

# Design, evaluation and future projections of the ~~NARClM2.0~~NARClM 2.0 CORDEX-CMIP6 Australasia regional climate ensemble

Giovanni Di Virgilio<sup>1,2</sup>, Jason P. Evans<sup>2,3</sup>, Fei Ji<sup>1,3</sup>, Eugene Tam<sup>1</sup>, Jatin Kala<sup>4</sup>, Julia Andrys<sup>4</sup>, Christopher Thomas<sup>2</sup>, Dipayan Choudhury<sup>1</sup>, Carlos Rocha<sup>1</sup>, Stephen White<sup>1</sup>, Yue Li<sup>1</sup>, Moutassem El Rafei<sup>1</sup>, Rishav Goyal<sup>1</sup>, Matthew L. Riley<sup>1</sup> and Jyothi Lingala<sup>4</sup>

<sup>1</sup>Climate & Atmospheric Science, NSW Department of Climate Change, Energy, the Environment and Water, Sydney, Australia

<sup>2</sup>Climate Change Research Centre, University of New South Wales, Sydney, Australia

<sup>3</sup>Australian Research Council Centre of Excellence for Climate Extremes, University of New South Wales, Sydney, Australia

<sup>4</sup>Environmental and Conservation Sciences, and Centre for Climate Impacted Terrestrial Ecosystems, Harry Butler Institute, Murdoch University, Murdoch, WA 6150, Australia

Correspondence to: Giovanni Di Virgilio ([giovanni.divirgilio@environment.nsw.gov.au](mailto:giovanni.divirgilio@environment.nsw.gov.au);  
[giovanni@unsw.edu.au](mailto:giovanni@unsw.edu.au))

1 **Abstract.** ~~NARClM2.0~~NARClM 2.0 comprises two Weather Research and Forecasting (WRF)  
2 regional climate models (RCMs) downscaling five CMIP6 global climate models contributing to the  
3 Coordinated Regional Downscaling Experiment over Australasia at 20 km resolution, and south-east  
4 Australia at 4 km convection-permitting resolution. We first describe ~~NARClM2.0~~NARClM 2.0's  
5 design, including selecting two, definitive RCMs via testing seventy-eight RCMs using different  
6 parameterisations for planetary boundary layer, microphysics, cumulus, radiation, and land surface  
7 model (LSM). We then assess ~~NARClM2.0~~NARClM 2.0's skill in simulating the historical climate  
8 versus CMIP3-forced ~~NARClM1.0~~NARClM 1.0 and CMIP5-forced ~~NARClM1.5~~NARClM 1.5  
9 RCMs and compare differences in future climate projections. RCMs using the new Noah-MP LSM in  
10 WRF with default settings confer substantial improvements in simulating temperature variables versus  
11 RCMs using Noah-Unified. Noah-MP confers smaller improvements in simulating precipitation,  
12 except for large improvements over Australia's southeast coast. Activating Noah-MP's dynamic  
13 vegetation cover and/or runoff options primarily improve simulation of minimum temperature.  
14 ~~NARClM2.0~~NARClM 2.0 confers large reductions in maximum temperature bias versus  
15 ~~NARClM1.0~~NARClM 1.0 and 1.5 (1.x), with small absolute biases of ~0.5K over many regions  
16 versus over ~2K for NARClM1.x. ~~NARClM2.0~~NARClM 2.0 reduces wet biases versus  
17 NARClM1.x by as much as 50%, but retains dry biases over Australia's north.

18 ~~NARCIIM2.0~~NARCIIM 2.0 is biased warmer for minimum temperature versus  
19 ~~NARCIIM1.5~~NARCIIM 1.5 which is partly inherited from stronger warm biases in CMIP6 versus  
20 CMIP5 GCMs. Under shared socioeconomic pathway (SSP)3-7.0, ~~NARCIIM2.0~~NARCIIM 2.0  
21 projects ~3K warming by 2060-79 over inland regions versus ~2.5K over coastal regions.  
22 ~~NARCIIM2.0~~NARCIIM 2.0-SSP3-7.0 projects dry futures over most of Australia, except for wet  
23 futures over Australia's north and parts of western Australia which are largest in summer.  
24 ~~NARCIIM2.0~~NARCIIM 2.0-SSP1-2.6 projects dry changes over Australia with only few exceptions.  
25 ~~NARCIIM2.0~~NARCIIM 2.0 is a valuable resource for assessing climate change impacts on societies  
26 and natural systems and informing resilience planning by reducing model biases versus earlier  
27 NARCIIM generations and providing more up-to-date future climate projections utilising CMIP6.

**Keywords:**

28 Climate change; climate impact adaptation; dynamical downscaling; CORDEX-CMIP6; model  
29 design; model evaluation

## 30 1. Introduction

31 Climate projections are foundational to informing climate change mitigation and adaptation planning  
32 at various spatial scales (IPCC, 2021). Regional climate models (RCMs) dynamically downscale  
33 global climate models (GCMs) at ~100-200 km resolution to simulate higher resolution climate  
34 projections that better resolve local-scale influences on regional climate, such as mountain ranges,  
35 land-use variation, land-sea contrasts, and convective processes (Torma et al., 2015; Giorgi, 2019). As  
36 such, whilst GCMs are the best tools for investigating climate at global scales, RCMs provide  
37 improved guidance for climate policy at regional scale, which is the scale at which climate change  
38 impacts are experienced (Hsiang et al., 2017).

39 The NARClIM programme (New South Wales and Australian Regional Climate Modelling) is  
40 now in its third generation. Like its predecessors, NARClIM version 2.0 ([‘NARClIM2.0NARClIM](#)  
41 [2.0](#)’), aims to produce robust, detailed regional climate projections at spatial scales relevant for use in  
42 local-scale climate change analysis. A key feature of all NARClIM generations is to simulate the  
43 climate over the Coordinated Regional Downscaling Experiment (CORDEX)-Australasia domain, and  
44 a higher resolution inner domain over southeast Australia via one-way nesting (Figure 1). With one-  
45 way nesting the inner domain obtains its initial and lateral boundary conditions from the simulation  
46 over CORDEX-Australasia. [NARClIM1.0NARClIM 1.0](#) simulated the climate of Australasia for three  
47 periods (1990-2009, 2020-2039, 2060-2079) at 50 km resolution and southeast Australia at 10 km  
48 using three configurations of the weather research and forecasting (WRF) RCM (Skamarock et al.,  
49 2008) to downscale GCMs from Coupled Model Intercomparison Project phase three (CMIP3) under  
50 the SRES A2 greenhouse gas (GHG) scenario (Evans et al., 2014). [NARClIM1.5NARClIM 1.5](#) used  
51 CMIP5 GCMs under representative concentration pathways (RCP) 4.5 and 8.5 to simulate  
52 continuously for 1950-2100 on the same grids as [NARClIM1.0NARClIM 1.0](#) using two of its RCMs  
53 (Nishant et al., 2021).

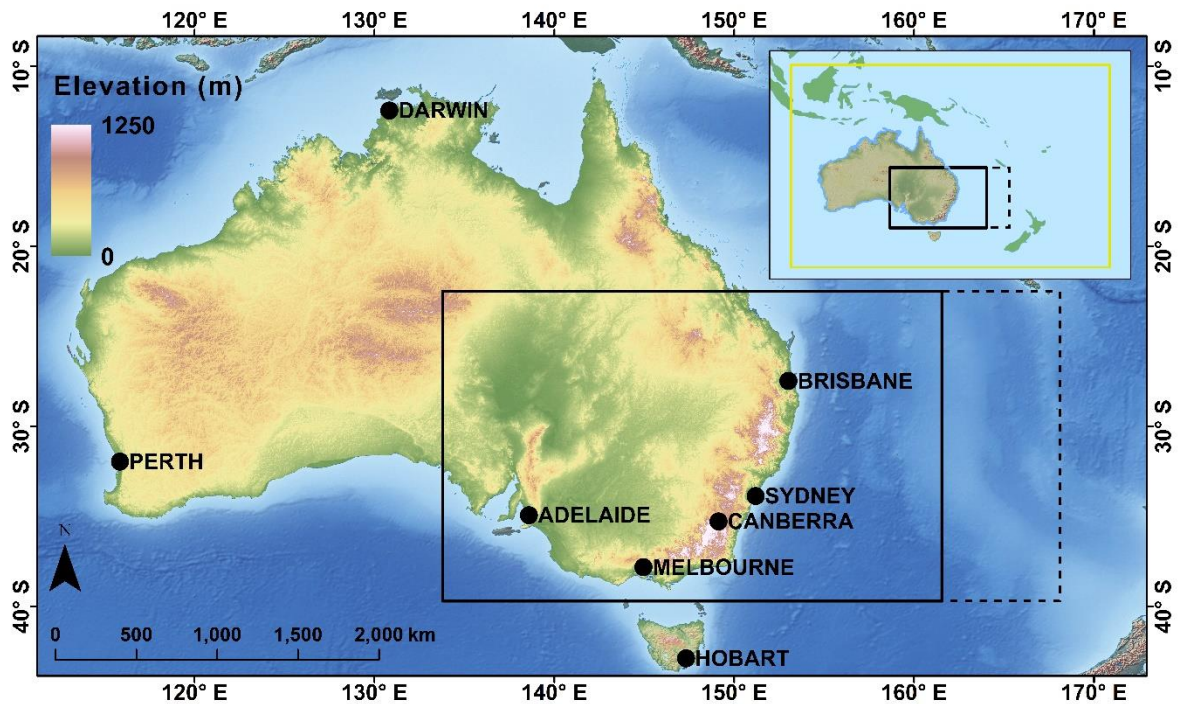
54 [NARClIM2.0NARClIM 2.0](#) aims to improve performance in simulating the Australian climate  
55 relative to previous NARClIM generations with the goal of better informing community resilience to  
56 climate change (New South Wales Government, 2022, 2023). All NARClIM projects include a  
57 bottom-up design ethos involving multi-sectoral end-user engagement in specifying model  
58 requirements to ensure model performance and outputs meet end-user needs. Key requirements from  
59 the [NARClIM2.0NARClIM 2.0](#) user-consultation include providing increased detail in climate  
60 simulations via higher resolution and improving the simulation of precipitation and temperature as  
61 these are fundamental inputs to climate impact studies. Whilst [NARClIM1.0NARClIM 1.0](#) and 1.5  
62 (1.x) confer the expected level of performance in simulating the Australian climate (Di Virgilio et al.,  
63 2019; Evans et al., 2020b), recent technological and scientific advancements mean that aspects of  
64 their performance might now be improved. NARClIM1.x RCMs show widespread cold biases in  
65 maximum temperature exceeding -5K for some RCMs. Conversely, minimum temperature is

66 simulated more accurately with biases in the range of  $\pm 1.5\text{K}$ . NARClIM1.x RCMs overestimate  
67 precipitation, particularly over Australia's socio-economically important eastern seaboard (Di Virgilio  
68 et al., 2019).

69 As they are expensive to run from both computational and data storage perspectives, dynamical  
70 downscaling projects like ~~NARClIM2.0~~NARClIM 2.0 use a subset of available GCMs as driving data,  
71 necessitating careful model selection. Similarly, a large combination of different physical  
72 parametrisations available for the WRF RCM enables many structurally different RCMs to be  
73 potentially used to downscale GCMs. A key component of ~~NARClIM2.0~~NARClIM 2.0's design is  
74 testing the viability of alternative RCM parameterisations via a three-phase approach, with each phase  
75 building on the preceding phase to identify the RCM parameterisations that perform well during  
76 testing to meet ~~NARClIM2.0~~NARClIM 2.0's aim of improving the simulation of Australia's climate.  
77 GCM and RCM statistical independence are also sought to avoid creating a biased sample of climate  
78 change. Hence, the aims of this paper are to:

- 79 1) describe how and why ~~NARClIM2.0~~NARClIM 2.0 differs from its predecessors in terms of  
80 its design and production processes, explaining the model test and evaluation approaches underlying  
81 its design decisions. A key focus is on the design and testing of seventy-eight structurally different  
82 WRF RCMs and their evaluation to identify a subset of RCMs for use in ~~NARClIM2.0~~NARClIM 2.0;  
83 2) characterise the performance improvements of CMIP6-~~NARClIM2.0~~NARClIM 2.0 RCMs in  
84 simulating the Australian climate relative to previous NARClIM generations by evaluating their skill  
85 in simulating mean maximum and minimum temperature and precipitation versus observations;  
86 ~~and~~ 3) summarise the climate projections produced by CMIP6-~~NARClIM2.0~~NARClIM 2.0 and  
87 how these differ from previous CMIP3-5-NARClIM generations.

88 The following section summarises the basic design features of each NARClIM generation;  
89 ~~section 3. describes NARClIM2.0's design process with a focus on its RCM physics testing, as well as~~  
90 ~~a brief overview of its production process;~~ sSection 4.3. describes evaluation methods and metrics;  
91 ~~Section 4. describes NARClIM 2.0's design process with a focus on its RCM physics testing, as well~~  
92 ~~as a brief overview of its production process;~~ Section 5. summarises the RCM physics test results;  
93 ~~s~~Section 6. evaluates the performance of all NARClIM models in simulating the recent Australian  
94 climate; ~~s~~Section 7. provides an overview of their future projections; and Section 8. discusses key  
95 results and summarises this paper.



96

97 **Figure 1.** Model domains for NARClM regional climate simulations. The southeast inner domain for  
 98 [NARClM2.0](#) is delineated with a solid black rectangle; the corresponding inner domain for  
 99 [NARClM1.0](#) and 1.5 is delineated with a dashed black line. The elevated terrain of the Australian  
 100 Alps which form part of the Great Dividing Range is in eastern Australia. Inset shows the CORDEX-Australasia  
 101 outer domain.

## 102 2. Three generations of NARClM: model overviews

103 The design of [NARClM1.0](#) is described in Evans et al. (2014); [NARClM1.5](#)  
 104 [1.5](#) used the same design approach but used CMIP5 rather than CMIP3 GCMs. All generations of  
 105 NARClM use different versions of the WRF model ([Skamarock et al., 2008](#)) to perform dynamical  
 106 downscaling of GCMs since the WRF model goes through regular updates. The southeast Australian  
 107 inner domain captures five of Australia's eight capital cities (Figure 1) and over 75% of the Australian  
 108 population (Australian Bureau Statistics, 2024). Additionally, the inner domain captures coastal  
 109 regions that are characterised by topographic complexity and land-use class variation. Regions east of  
 110 the Great Dividing Range mountains in southeast Australia (Figure 1) show different responses to  
 111 oceanic climate modes compared to inland semi-arid regions (Murphy and Timbal, 2008) and are  
 112 impacted by events such as rapidly developing storms, including east coast lows (Pepler and Dowdy,  
 113 2021). Such atmospheric processes are not adequately resolved by GCMs due to coarse resolutions  
 114 (Di Virgilio et al., 2022; Grose et al., 2020).

115 [NARClM2.0](#) encompasses several design advancements over its predecessors  
 116 (Table 1). [NARClM2.0](#) RCMs have a 20 km resolution CORDEX-Australasia domain

117 (versus 50 km) and 4 km (versus 10 km) domain over southeast Australia and use 45 (versus 30)  
118 vertical levels. The aim of increasing the resolution of this inner domain from 10 km to 4 km is to  
119 render these simulations convection-permitting (Kendon et al., 2021; Lucas-Picher et al., 2021).  
120 Hence, whilst the 20 km-resolution outer domain uses cumulus parametrisation, simulations over the  
121 4 km domain do not use cumulus parametrisation. ~~NARClIM2.0~~NARClIM 2.0 also includes a new  
122 collaboration with the Western Australian government, with separate 4 km simulations being  
123 performed over south-west and north-west Western Australia (not shown in Figure 1) as part of the  
124 Western Australian climate science initiative (DWER, 2023). Boundary conditions derived from the  
125 20 km ~~NARClIM2.0~~NARClIM 2.0 CORDEX Australasia domain are used to drive these simulations.  
126 Additional major differences in model setup for ~~NARClIM2.0~~NARClIM 2.0 include:

- 127     ▪ ~~NARClIM1.0~~NARClIM 1.0 RCMs use different parameterisations for planetary boundary  
128     layer (PBL) physics, surface physics, cumulus physics, land surface model (LSM), and radia-  
129     tion (Evans et al., 2014). These RCM parameterisations were also used for ~~NAR-~~  
130     ~~ClIM1.5~~NARClIM 1.5. Owing to the project aims stated above, RCM parameterisations for  
131     ~~NARClIM2.0~~NARClIM 2.0 differ to those of NARClIM1.x (see [sSect. 34](#)).
- 132     ▪ ~~NARClIM2.0~~NARClIM 2.0 increases the number of driving GCMs to 5 and simulates for a  
133     wider range of plausible future climates via three shared socioeconomic pathways (SSP).  
134     SSP1-2.6 is selected as a low GHG scenario envisaging a future climate with CO<sub>2</sub> emissions  
135     cut to net zero by around 2075 and warming held to below 2°C by 2100; SSP2-4.5 estimates  
136     projected warming under a ‘middle of the road’ scenario where temperatures increase to  
137     ~2.7°C by 2100; and SSP3-7.0 is a high GHG scenario which assumes warming of ~4°C by  
138     2100 (IPCC, 2021).
- 139     ▪ Urban physics is activated in ~~NARClIM2.0~~NARClIM 2.0 (WRF setting: sf\_urban\_physics=1)  
140     to represent surface energy balance in urban areas via a single layer urban canopy model  
141     (Kusaka and Kimura, 2004).
- 142     ▪ Input of different aerosol species is activated for the RCM radiation scheme using the Tegen  
143     et al. (1997) climatology available in WRF (aer\_opt=1). This aerosol forcing is the same for  
144     all GCMs, and not model-specific.
- 145     ▪ The eastern boundary of the ~~NARClIM2.0~~NARClIM 2.0 inner domain is located further  
146     westward relative to that of NARClIM1.x (Figure 1).

147 **Table 1.** High-level design features of three generations of NARCLiM regional climate models

	<b>Model Generation</b>		
	<del>NARCLiM 1.0</del> <u>NARCLiM 1.0</u>	<del>NARCLiM 1.5</del> <u>NARCLiM 1.5</u>	<del>NARCLiM 2.0</del> <u>NARCLiM 2.0</u>
<b>Release date</b>	2014	2020	2023-2024
<b>Years simulated</b>	1990-2009, 2020-2039, 2060-2079	1950-2100	1950-2100
<b>Grid resolutions: CORDEX-Australasia; NARCLiM inner domains</b>	50 km; 10 km	50 km; 10 km	20 km; 4 km
<b>Vertical levels</b>	30	30	45
<b>Global Climate Models</b>	4 CMIP3 GCMs	3 CMIP5 GCMs	5 CMIP6 GCMs
<b>Regional Climate Models</b>	3 RCM configurations (WRF3.3)	2 RCM configurations (WRF3.6.0.5)	2 RCM configurations (WRF4.1.2)
<b>Future emission scenarios</b>	SRES A2	RCP4.5, RCP8.5	SSP1-2.6, SSP2-4.5, SSP3-7.0
<b>Reanalysis-driven (CORDEX Evaluation)</b>	NCEP: 1950-2009	ERA-Interim: 1979-2013	ERA5: 1979-2020
<b><u>Computational resources</u> (core hours)</b>	<u>30M</u>	<u>30M</u>	<u>1060M</u>

148 **3. Evaluation methods**

149 This section largely focuses on the methods and metrics used for the NARCLiM 2.0 RCM physics test-  
150 ing and comparisons of model biases and future climate projections against previous generations of  
151 NARCLiM. Details on methods and results for the CMIP6 GCM evaluation used to select driving  
152 GCMs and the ERA5-NARCLiM 2.0 RCM evaluation used to select two, definitive RCMs for the  
153 GCM-driven simulations are available in Di Virgilio et al. (2022) and Di Virgilio et al. (in review),  
154 respectively, with overviews of these components of NARCLiM 2.0 design provided in Sections 4.2  
155 and 4.4 below.

### 156 **3.1 Observations**

157 Australian Gridded Climate Data (AGCD version 1.0; (Evans et al., 2020a) are the observational data  
158 used to evaluate the NARClIM 2.0 RCM physics test RCMs. These daily gridded data for maximum  
159 and minimum temperature and precipitation are obtained from an interpolation of station observations  
160 across Australia. AGCD data are on a regular WGS84 grid with a grid-averaged resolution of 0.05°.  
161 For the NARClIM 2.0 RCM physics tests, the AGCD data were re-gridded to correspond with the  
162 RCM data from the inner domain on their native grids using a conservative area-weighted re-gridding  
163 scheme. All data (RCM and AGCD) were restricted to a common extent contained within the inner  
164 domain over southeast Australia, and a land mask was applied so that statistics were computed using  
165 only land pixels. Treatment of AGCD for the CMIP6 GCM evaluation and the ERA5-NARClIM 2.0  
166 RCM evaluation is described in Di Virgilio et al. (2022) and Di Virgilio et al. (in review), respective-  
167 ly.

### 168 **3.2 Methods and metrics: phase I-III NARClIM2.0 physics tests**

169 Test RCM performances in reproducing observations for daily maximum and minimum temperature  
170 and daily precipitation were assessed by calculating the model bias, i.e., model outputs minus AGCD,  
171 and the RMSE of modelled versus observed fields. Model biases and RMSEs were calculated at an-  
172 annual and seasonal timescales. The model representations of the hottest and the wettest day on an an-  
173 annual time scale over the study region were also compared with AGCD. Probability density functions  
174 (PDFs) were calculated for each variable using daily data. The Perkins skill score (PSS) (Perkins et  
175 al., 2007) was calculated to assess the overall degree of overlap between modelled and observed dis-  
176 tributions, with PSS = 1 indicating that distributions overlap perfectly.

177 There are several methods to evaluate the overall performance of RCMs. In this study, we  
178 ranked the RCMs individually based on their bias, RMSE, and PSS for maximum temperature, mini-  
179 um temperature, and precipitation. Each variable was ranked separately for each metric. The ranks  
180 were then summed to determine the overall ranking for each RCM.

### 181 **3.3 Independence assessments**

182 We used the method of Bishop and Abramowitz (2013) as one of two methods of assessing the inde-  
183 pendence of physics test RCMs and the target CMIP6 GCMs under evaluation for use in NARClIM  
184 2.0. This approach uses the covariance in model errors as the basis to define model dependence; spe-  
185 cifically, independence coefficients are derived from the error covariance matrix of the RCMs or  
186 GCMs. Model independence is quantified using the correlation of model errors. For the physics test  
187 RCMs, errors are computed by comparing the climatology of maximum and minimum temperature  
188 and precipitation over the south-east Australia inner domain for 2016 with corresponding AGCD ob-



189 servations. The same calculation is performed for the CMIP6 GCMs, except for the Australian conti-  
190 nent. Daily timeseries of precipitation, maximum and minimum temperature are calculated individual-  
191 ly for each RCM and for AGCD. The simulated and observed daily timeseries of each variable are  
192 then normalised by the standard deviation of the corresponding observed variable. These normalised  
193 variables are concatenated for each RCM (GCM) and AGCD. An anomaly time series for each grid  
194 cell is then produced. These time series are used to create a model error covariance matrix containing  
195 the errors for all RCMs (GCMs). The coefficients of a linear combination of the RCMs (GCMs) that  
196 optimally minimises the mean square error depends on both model performance and model depend-  
197 ence (Bishop and Abramowitz, 2013). The result of this minimisation problem is written in terms of  
198 the covariance matrix. The magnitude of coefficients assigned to each RCM (GCM) reflects a combi-  
199 nation of their performance and independence. Highly independent models have different errors when  
200 simulating the recent climate. Models with the largest coefficients have the most independent errors  
201 versus observations.

202 \_\_\_\_\_ The Herger method of subset selection (Herger et al., 2018), as implemented here, uses quad-  
203 ratic integer programming to find the subset of models whose equally-weighted subset mean (EWSM)  
204 minimises a quadratic cost function. This cost function is chosen to measure the performance of the  
205 EWSM in comparison to a given observational product. The two cost functions used here are: the  
206 mean squared error (MSE) between the EWSM and the observational product (Herger et al. 2018, Eq.  
207 1); and another which measures a combination of the MSE of the EWSM, the average MSE of each  
208 subset member, and the average pairwise mean squared distance between subset members (Herger et  
209 al. 2018, Eq. 2).

### 210 **3.4 NARcliM2 CMIP6-RCMs: historical evaluation and climate change** 211 **projections**

212 Performances of NARcliM 2.0 versus NARcliM1.x RCMs in reproducing the recent Australian cli-  
213 mate are evaluated by calculating the model biases (model outputs minus AGCD observations) for  
214 mean maximum and minimum temperature and precipitation for 1990-2009. To enable comparison of  
215 future projections between NARcliM 1.0, NARcliM 1.5 and NARcliM 2.0 (where NARcliM 1.0  
216 modelled for 1990-2009, 2020-2039, and 2060-2079), all NARcliM ensemble projected changes are  
217 shown as far future (2060–2079) minus present day (1990–2009).

### 218 **3.5 Statistical significance**

219 When quantifying RCMs' future climate change projections (compared to the historical period) and  
220 biases in maximum and minimum temperature, the statistical significance is calculated for each grid  
221 cell using t-tests assuming equal variance. The Mann–Whitney U test is used for precipitation given  
222 its non-normality. Significance thresholds were adjusted to account for multiple testing using Walk-

er's test (Eq.2 in Wilks, 2016). For individual RCMs, grid cells showing statistically significant changes are stippled, otherwise they are shown in colour where change is statistically insignificant. Results on the statistical significance of each ensemble mean are separated into three categories following Tebaldi et al. (2011): 1) statistically insignificant areas are shown in colour, denoting that less than 50% of RCMs are significantly biased/different; 2) in areas of significant agreement (stippled), at least 50% of RCMs are significantly biased/different and at least 70% of significant models in the CMIP6-NARClIM 2.0 RCM ensemble agree on the sign of the bias/difference. In such areas, many ensemble members have the same bias sign which is an undesirable outcome; and 3) areas of significant disagreement, where at least 50% of RCMs are significantly biased/different and fewer than 70% of significant models agree on the bias sign, are shown with diagonal hatching for the CMIP6-NARClIM 2.0 historical evaluation and climate change signals.

## 34. ~~NARClIM2.0~~NARClIM 2.0 design and production process

### overview

The ~~NARClIM2.0~~NARClIM 2.0 design and production processes are summarised below in reference to Figure 2. The design process is an adaptation of that introduced in Evans et al. (2014). Two companion manuscripts describe elements shown in Figure 2, ~~and~~ which are therefore only summarised briefly in this manuscript: Di Virgilio et al. (2022) describes the CMIP6 GCM selection process summarised in Box 2, and Di Virgilio et al. (in review) describes the ERA5-~~RCM~~ evaluation undertaken in Boxes 5 and 6.

#### I. Design Phase:

- i) **Box 1:** model design requirements are identified via consultation between ~~NARClIM2.0~~NARClIM 2.0 modelling groups and multi-sectoral end-users, as well as adherence to CORDEX-CMIP6 design requirements (W~~CRP~~etp, 2020).
- ii) **Box 2:** NARClIM1.x selected driving CMIP3-5 GCMs (respectively) via literature review of existing GCM evaluations. During ~~NARClIM2.0~~NARClIM 2.0 design, there were no pre-existing comprehensive evaluations of individual CMIP6 GCMs for the Australian region, including assessments of climate change signals and GCM statistical independence. Hence, an evaluation and selection of CMIP6 GCMs was conducted (see Di Virgilio et al. 2022). This evaluation selected five GCMs to force two ~~NARClIM2.0~~NARClIM 2.0 RCMs (see ~~S~~sect 4.2 and 4.4). The relative contribution to uncertainty/variation in climate projections can be larger for GCMs than for RCMs (e.g. Lee et al., 2023).
- iii) **Boxes 3-4:** a new WRF RCM multi-physics test ensemble is created for ~~NARClIM2.0~~NARClIM 2.0: RCM physics testing is conducted via a three-phase approach, with each phase building on the findings of the preceding phase to identify the RCM pa-

257 parameterisations that perform well during testing with the aim of improving the simulation  
258 of the Australian climate. In this way, RCMs are parameterised with different physics set-  
259 tings via each test phase, systematically removing poor performing options while facilitat-  
260 ing the fine tuning and improvement of the parameterisations that perform well during  
261 testing to build a total ensemble size of seventy-eight structurally different test RCMs.  
262 The performances of the different test RCM configurations ~~are~~ is evaluated, ultimately  
263 leading to the selection of a subset of seven RCMs for subsequent downscaling of ERA5  
264 reanalysis ~~and comprising~~ as part of the CORDEX evaluation experiment.

265 iv) **Boxes 45-6:** These seven RCMs are used to downscale ERA5 reanalysis over the 20 km  
266 and 4 km domains for 1979-2020. Evaluating these ERA5-forced simulations informs se-  
267 lection of two definitive, <sup>1</sup>production<sup>2</sup> RCMs for CMIP6-forced downscaling (see sSect.  
268 4.4 and Di Virgilio et al. in review).

## 269 II. Production Phase:

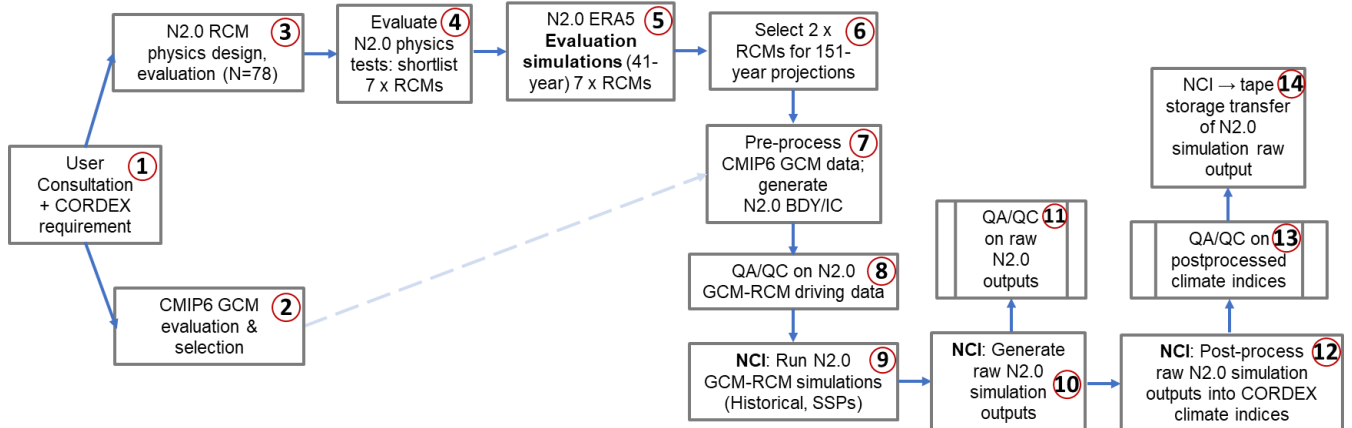
- 270 i) **Boxes 7-8:** CMIP6 GCM data are pre-processed to create initial and boundary conditions  
271 to drive simulations for the historical (1950-2014) and SSP experiments (2015-2100). A  
272 code repository used for this GCM preprocessing is available on Zenodo at:  
273 <https://doi.org/10.5281/zenodo.11184830>~~[https://bitbucket.org/oeheas/narelim2-](https://bitbucket.org/oeheas/narelim2-0-design-and-evaluation-2024-support-materials/src/main/)~~  
274 ~~[0-design-and-evaluation-2024-support-materials/src/main/](https://bitbucket.org/oeheas/narelim2-0-design-and-evaluation-2024-support-materials/src/main/)~~ within the  
275 WRF/repo\_snapshots subdirectory. Quality assurance/quality control (QA/QC) is per-  
276 formed on these data before initiating the simulations (e.g. variables are checked to con-  
277 firm data do not contain significant outliers across ensemble members).
- 278 ii) **Boxes 9-11:** the 151-year CMIP6-forced ~~NARCIIM2.0~~NARCIIM 2.0 RCM simulations  
279 are run using National Computational Infrastructure at Canberra, Australia (NCI,  
280 <https://nci.org.au/>). File integrity verification and QA/QC are performed on each year of  
281 raw WRF output throughout the simulation lifecycle and prior to post-processing to  
282 CORDEX-compliant format climate variables. QA/QC tests include calculating the min-  
283 imum, maximum, mean and standard deviation for key variables over consecutive periods  
284 of six simulation days. Variables are categorised as either normally distributed or other-  
285 wise. Normally distributed variables (e.g. surface temperature) are deemed potentially er-  
286 roneous if their minima/maxima are greater than five standard deviations away from the  
287 global mean of the relevant statistic of the rolling six-day period. Non-normally distribut-  
288 ed variables (e.g. snow depth and precipitation) are checked only for global minima and  
289 maxima ~~only~~.
- 290 iii) **Boxes 12-13:** after each year of simulation raw output is generated, their post-processing  
291 is initiated to produce CORDEX CORE, Tier 1 and Tier 2 variables (WCRP, 2022). A  
292 statistical QA/QC process is automatically applied to each year of post-processed

293 CORDEX CORE variables as they are generated throughout the simulations. QA/QC  
 294 tests include:

- 295     ▪ Check for presence of missing values.
- 296     ▪ Check that all values are within realistic ranges for minima and maxima.
- 297     ▪ Check minima and maxima are not equal at any timestep with exceptions (e.g.,  
 298       snow depth which can be zero everywhere in the outer domain).
- 299     ▪ Check that changes over time are within realistic ranges (~~i.e.~~, assess temporal  
 300       gradients).
- 301     ▪ Check that changes between neighbouring data points are within realistic ranges  
 302       (~~i.e.~~, assess spatial gradients).
- 303     ▪ Check the number of grid cells with NaN (non-numerical) values do not exceed  
 304       the threshold set for the variable.

305 Reasonable ranges for variables are determined using a series of threshold values that are  
 306 based on historical records and/or empirical analysis. QA/QC computer scripts generate  
 307 'exceedance files' which output every data point that surpasses the threshold values, and  
 308 these exceedance files are then manually reviewed to determine whether an issue is a true  
 309 or false positive, etc.

310 iv) **Box 14:** Once each year of WRF raw files ~~are~~ is post-processed, raw files are transferred  
 311 to a tape facility for long-term storage.



312  
 313 **Figure 2.** Simplified overview of ~~NARCI M2.0~~ NARCI M 2.0 (N2.0) design and production processes. ERA5 =  
 314 ECMWF Reanalysis v5 data; BDY = boundary conditions; IC = Initial conditions; QA/QC = Quality Assurance  
 315 / Quality Control; NCI = National Computationalg Infrastructure (high performance computer used ~~for~~ to run  
 316 N2.0 ~~production~~ simulations).

317 These model design and production stages are now described in more detail:

## 318 **43.1 Model evaluation and selection**

319 Practical constraints such as available compute and data storage resources enforce an upper limit on  
320 GCM-RCM ensemble size. Thus, ~~NARCIIM2.0~~NARCIIM 2.0 uses a subset of available CMIP6  
321 GCMs and WRF RCM configurations, necessitating careful GCM and RCM selection to create a  
322 subset of GCM-RCMs that provide robust climate simulations whilst also adequately sampling model  
323 uncertainty. In selecting a subset of GCMs and RCMs for dynamical downscaling, it is desirable to  
324 reject models that perform consistently poorly relative to their peers in simulating the current climate,  
325 as this provides lower confidence in the projected change (Evans et al., 2020b; Di Virgilio et al.,  
326 2022; Grose et al., 2023). Furthermore, the modelled climate space sampled is reduced when selecting  
327 a subset of GCMs, which can create a biased view of the climate, as well as the plausible change in  
328 climate. Care must therefore be taken to ensure that the subset of models used for downscaling are  
329 representative of the full range of possible climates, and that model errors are uncorrelated, ~~i.e.~~i.e.,  
330 that models are statistically independent. The steps taken to evaluate and select GCMs and RCMs for  
331 ~~NARCIIM2.0~~NARCIIM 2.0 are described next.

## 332 **43.2 CMIP6 GCM evaluation**

333 A three-phase process was used to evaluate individual CMIP6 GCMs (for further details see Di  
334 Virgilio et al. 2022):

### 335 **43.2.1 CMIP6 GCM Performance**

336 ~~The~~We evaluated the performances of individual CMIP6 GCMs ~~in simulating the Australian climate~~  
337 ~~were assessed with respect to climate means, extremes, climate modes, and daily climate variable~~  
338 ~~distributions their skill~~ in simulating the following aspects of the observed historical climate of  
339 Australia:

- 340 ▪ annual and seasonal climatologies and daily distributions of maximum and minimum temper-  
341 atures and precipitation.
- 342 ▪ climate extremes, such as the 99<sup>th</sup> percentiles of daily maximum temperature and precipita-  
343 tion, and the 1<sup>st</sup> percentile of minimum temperature.
- 344 ▪ teleconnections of oceanic climate modes and Australian regional rainfall.

345 Temperature and precipitation variables are chosen for evaluation because, being well-represented in  
346 high-quality gridded observational data sets for the Australian continent, they provide the most direct  
347 comparison to observations (King et al., 2013). They are also often prioritised for impact studies.  
348 Given variables such as winds (U, V), air temperature (T), water mixing ratio (Q), geopotential height  
349 (Z), sea surface temperature (SST), and sea level pressure (PSL) serve as boundary conditions for

350 driving RCMs, these could be incorporated into future GCM evaluation studies. However, evaluating  
351 such variables would require use of re-analysis data as surrogate observations.

352 A set of GCMs that performed consistently poorly across the variables and statistics  
353 considered were identified. These models, as well as those with insufficient data to enable dynamical  
354 downscaling using the WRF RCM, were excluded from further evaluation leaving 27 GCMs for  
355 subsequent assessment.

### 356 **4.3.2.2 CMIP6 GCM Independence**

357 The retained 27 GCMs were subjected to the Bishop and Abramowitz (2013) and Herger et al. (2018)  
358 independence analyses (see [Sect. 3.5](#)). The GCMs were then ranked according to their relative level  
359 of statistical independence.

### 360 **4.3.2.3 Sampling CMIP6 GCM Climate Change Spread**

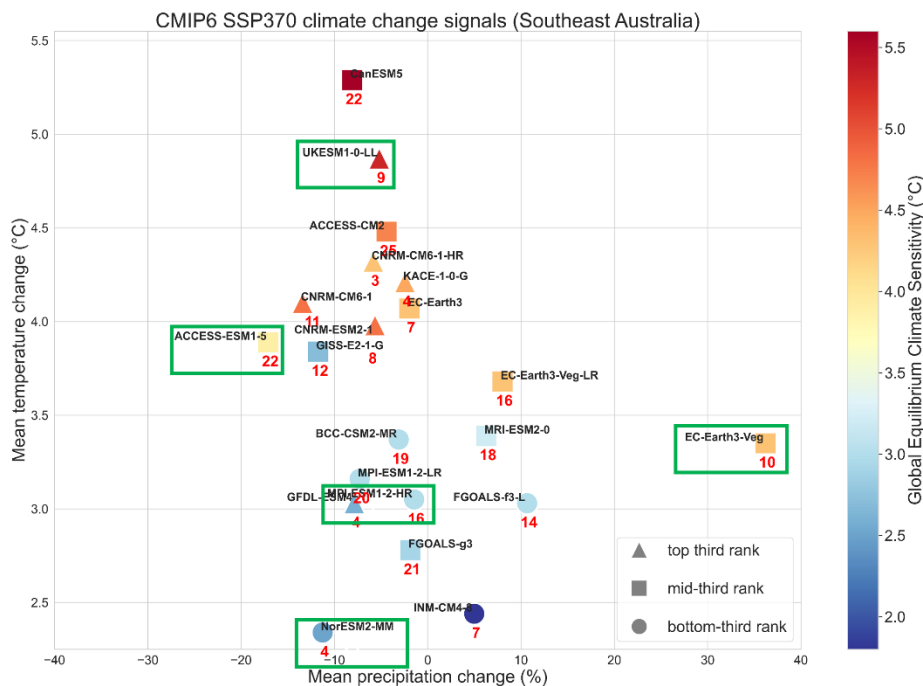
361 For climate change risk assessments, climate projections should reflect as much of the range of  
362 plausible future climate changes as possible (Whetton and Hennessy, 2010). The subset of CMIP6  
363 GCMs selected for ~~NARCM2.0~~[NARCM 2.0](#) spanned a wide range of future changes in annual  
364 mean temperature and precipitation. Climate change signals were calculated for 2080-2099 minus  
365 1995-2014 for the Australian continent and south-east Australia under SSP3-7.0 (for the latter, see  
366 Figure 3). The GCM independence rankings were placed within this climate change space, with  
367 higher independence rankings viewed as favourable, along with consideration of the following  
368 criteria:

- 369 i) A balanced range of GCM Equilibrium Climate Sensitivities (ECS) were sampled. ECS is the  
370 long-term increase in global mean surface air temperature in response to the radiative forcing  
371 caused by a doubling of pre-industrial CO<sub>2</sub> concentrations. ECS is related to global tempera-  
372 ture change, not just changes over Australia, however, it correlates strongly with regional  
373 warming. Around one third of CMIP6 GCMs show ECS values higher than the upper end of  
374 the likely range of 2.5°C to 4°C (IPCC, 2021). An upper range of > ~5°C cannot be ruled out  
375 (Meehl et al., 2020; Bjordal et al., 2020; Sherwood et al., 2020).
- 376 ii) Some CMIP6 GCMs that are favourable in terms of model performance and independence  
377 could not be selected as input to WRF for ~~NARCM2.0~~[NARCM 2.0](#) owing to insufficient  
378 data availability for key variables/~~variable~~, where ideally, WRF requires sub-daily data for the  
379 variables shown in Supporting Information, Table S1.

380 As a result of the above process, the five CMIP6 GCMs listed in Table 2 are selected to force each of  
381 the two definitive ~~NARCM2.0~~[NARCM 2.0](#) RCMs selected via the RCM physics testing and ERA5  
382 evaluation processes.

383 **Table 2.** Basic details of the CMIP6 GCMs used to force the two definitive NARCIIM2.0 simula-  
 384 tions RCMs comprising the NARCIIM 2.0 CORDEX-CMIP6 ensemble.

CMIP6 GCM	Institution	Variant/Run	Atmosphere lat/lon grid (°)
ACCESS-ESM1-5	CSIRO	r6i1p1f1	1.2 × 1.8
EC-Earth3-Veg	EC-EARTH consortium	r1i1p1f1	0.7 × 0.7
MPI-ESM1-2-HR	Max Planck Institute for Meteorology (MPI)	r1i1p1f1	~0.9
NorESM2-MM	Norwegian Climate Centre	r1i1p1f1	0.9 × 0.9
UKESM1-0-LL	UK Met Office and NERC research centres	r1i1p1f2	1.3 × 1.9



385 **Figure 3.** CMIP6 GCM climate change signals (2080-2099 versus 1995-2014) over south-east Australia for the  
 386 subset of GCMs retained following the model performance evaluation in Di Virgilio et al. (2022), and that  
 387 simulated at least monthly mean near surface air temperature and precipitation for the SSP-3.70 scenario. Boxed  
 388 GCMs are selected to force NARCIIM2.0 RCMs. Marker shapes indicate overall GCM  
 389 performance; markers are coloured according to their global equilibrium climate sensitivity (ECS) values; **Red**  
 390 numbers represent the smallest Herger Method 1 set for that GCM.  
 391

### 392 43.3 NARCIIM 2.0 RCM physics testing

393 The NARCIIM 2.0 RCM physics testing aims to identify and exclude RCMs that  
 394 perform consistently poorly in simulating the southeast Australian climate and to select RCMs that  
 395 have high statistical independence. The selection of RCMs in NARCIIM 2.0 involves  
 396 the creation of a multi-physics ensemble where each RCM uses different physical parametrisations for  
 397 PBL, microphysics, cumulus, radiation, and LSM. This enables many structurally different RCMs to  
 398 be constructed and tested. In NARCIIM 1.0, 36 WRF RCM configurations were

399 designed, tested, and evaluated (Evans et al. 2014). ~~NARCIIM2.0~~NARCIIM 2.0 physics testing  
400 assesses 78 RCM configurations which are progressively tested via three phases, where each test  
401 phase is informed by the outcomes of the preceding phase to systematically remove poor performing  
402 RCM options while facilitating the selection of parameterisations that perform well during testing.  
403 The N=36 RCMs tested for ~~NARCIIM1.0~~NARCIIM 1.0 were evaluated based on eight representative  
404 storm event simulations each of two-weeks duration (Evans et al. 2014). ~~NARCIIM2.0~~NARCIIM 2.0  
405 physics simulations were run over an entire annual cycle (2016) with a two-month spin-up period  
406 commencing 1 November 2015. Australia experienced a range of weather extremes during 2016  
407 driven by a range of climatic influences making 2016 a suitable target year (Bureau of Meteorology,  
408 2017). Whilst assessing RCMs for an entire year improves on assessing for discrete storm events as  
409 per physics testing for ~~NARCIIM1.0~~NARCIIM 1.0, it was not feasible to run a large RCM physics  
410 ensemble for a longer duration. Initial and boundary conditions for all phases of the  
411 ~~NARCIIM2.0~~NARCIIM 2.0 RCM physics test simulations were derived from the ERA-Interim  
412 reanalysis data set (Dee et al., 2011). ERA-Interim was used because ERA5 was not available at the  
413 time. The three phases of ~~NARCIIM2.0~~NARCIIM 2.0 physics testing are as follows:

#### 414 **4.3.3.1 Phase I (N=36)**

415 Thirty-six RCMs ~~were~~ evaluated in Phase I. One radiation scheme (RRTMG) ~~was~~ tested for both  
416 long and short-wave radiation (it ~~was~~ held fixed for all RCMs), whereas physics settings for PBL,  
417 microphysics, cumulus, and LSM ~~were~~ varied. Of the 36 simulations, 18 used the Noah-Unified  
418 LSM, whilst the remainder used Community Land Model version 4.0 (CLM4). The physics options  
419 tested are listed in Table 3, where these were selected based on literature review. Each physics test  
420 simulation is denoted by a 12-digit identifier which comprises 6 pairs of digits, with each pair  
421 corresponding to the choice of a specific physics option as specified in the WRF namelist.input file.  
422 These pairs of digits follow the order: planetary boundary layer (pbl) | cloud microphysics (mp) |  
423 cumulus convection (cu) | shortwave radiation (sw) | longwave radiation (lw) | LSM (sf) and  
424 correspond to the WRF namelist options shown in Table 3. For example, the simulation  
425 ~~'050601040402'~~ is interpreted as: 05 | 06 | 01 | 04 | 04 | 02 and denotes that this simulation uses the  
426 following physics settings:

bl_pbl_physics	= 05 (MYNN2)
mp_physics	= 06 (WSM6)
cu_physics	= 01 (Kain-Fritsch)
ra_sw_physics	= 04 (RRTMG)
ra_lw_physics	= 04 (RRTMG)
sf_surface_physics	= 02 (Noah Unified)



427 The complete set of WRF RCM configurations tested in Phase I is shown in Supporting Information  
 428 Table S2.

429 **Table 3.** Physics options used in phase I (N=36) tests.

<u>Physics Option Description</u>	<u>WRF Namelist</u>	<u>Options Tested</u>	<u>Reference</u>
<u>Planetary boundary layer</u>	<u>bl_pbl_physics</u>	<u>01 = YSU</u>	Hong et al. (2006)
		<u>05 = MYNN2</u>	Nakanishi & Niino (2009)
		<u>07 = ACM2</u>	Pleim (2007)
<u>Microphysics</u>	<u>mp_physics</u>	<u>06 = WSM6</u>	Hong and Lim (2006)
		<u>08 = Thompson</u>	Thompson et al. (2008)
<u>Cumulus parameterisation</u>	<u>cu_physics</u>	<u>01 = Kain-Fritsch</u>	Kain (2004)
		<u>02 = BMJ</u>	Janjić (2000)
		<u>06 = Tiedtke</u>	Tiedtke (1989)
<u>Shortwave radiation</u>	<u>ra_sw_physics</u>	<u>04 = RRTMG</u>	Iacono et al. (2008)
<u>Longwave radiation</u>	<u>ra_lw_physics</u>	<u>04 = RRTMG</u>	-
<u>Land surface model</u>	<u>sf_surface_physics</u>	<u>02 = Noah-Unified</u>	Tewari et al. (2016)
		<u>05 = Community Land Model V4</u>	Oleson et al. (2010)

430

<u>Physics Option Description</u>	<u>WRF Namelist</u>	<u>Options Tested</u>
<u>Planetary boundary layer</u>	<u>bl_pbl_physics</u>	<u>01 = YSU</u>
		<u>05 = MYNN2</u>
		<u>07 = ACM2</u>
<u>Microphysics</u>	<u>mp_physics</u>	<u>06 = WSM6</u>
		<u>08 = Thompson</u>
<u>Cumulus parameterisation</u>	<u>cu_physics</u>	<u>01 = Kain-Fritsch</u>
		<u>02 = BMJ</u>
		<u>06 = Tiedtke</u>
<u>Shortwave radiation</u>	<u>ra_sw_physics</u>	<u>04 = RRTMG</u>
<u>Longwave radiation</u>	<u>ra_lw_physics</u>	<u>04 = RRTMG</u>
<u>Land surface model</u>	<u>sf_surface_physics</u>	<u>02 = Noah-Unified</u>
		<u>05 = Community Land Model V4</u>

431 **34.3.2 Phase II (N=60): additional LSM and radiation scheme tests**

432 Phase I RCMs using CLM4.0 were omitted from further testing because they did not consistently im-  
 433 prove performance in simulating the Australian climate relative to RCMs using Noah-Unified. In ad-  
 434 dition, RCMs using CLM4.0 had increased simulation times (by approximately twice when compared  
 435 to Noah-Unified). Hence, Phase II focuses exclusively on further testing of the RCM configurations  
 436 that used the Noah-Unified LSM.

437 The physics settings tested in Phase II are an alternative LSM to Noah-Unified (Noah Multi-  
 438 Parameterisation; 'Noah-MP2', Niu et al., 2011) and New Goddard radiation (Chou et al., 2001). Ow-

439 ing to time/resource constraints, testing all eighteen Phase I RCMs using Noah-Unified was not feasi-  
 440 ble. To reduce the number of RCMs for further testing, the worst-performing Noah-Unified based  
 441 RCM configurations identified in Phase I were excluded. The N=18 RCMs using Noah-Unified are  
 442 listed along with their overall performance total scores in Table 4 where the lowest scores under  
 443 ‘Rank totals’ indicate the RCMs that overall perform relatively well versus their peers (see sSect. 3  
 444 Evaluation Methods). Note that the ‘Overall rank’ denotes the RCMs’ relative ranking among all  
 445 Phase I RCMs. There is a sharp reduction in rank totals for RCMs #13-18 inclusive, relative to RCMs  
 446 #1-12. Therefore, RCMs #13-18 are excluded from further testing, and RCMs #1-12 are retained.

447 **Table 4.** RCM physics combination ranks of the Phase I, N=18 Noah Unified (NU) based RCMs.  
 448 Scores/ranks are based on model bias and root mean square error for annual and seasonal precipita-  
 449 tion, minimum temperature, maximum temperature, climate extremes (wettest and hottest days), and  
 450 Perkins Skill Scores (see sSect. 3). RCMs #1-12 are selected for further testing.

RCM #	RCM ID	Physics combination					Rank total	Overall rank in N=36 Phase I
		PBL	MP	Cumulus	SW/LW	LSM		
1	070801040402	ACM2	Thom	KF	RRTMG	NU	484	1
2	070601040402	ACM3	WSM6	KF	RRTMG	NU	495	2
3	070802040402	ACM4	Thom	BMJ	RRTMG	NU	527	3
4	070602040402	ACM5	WSM6	BMJ	RRTMG	NU	559	4
5	010802040402	YSU	Thom	BMJ	RRTMG	NU	574	7
6	050801040402	MYNN2	Thom	KF	RRTMG	NU	583	8
7	010801040402	YSU	Thompson	KF	RRTMG	NU	617	11
8	050802040402	MYNN2	Thompson	BMJ	RRTMG	NU	630	12
9	070606040402	ACM2	WSM6	Tiedtke	RRTMG	NU	639	13
10	050601040402	MYNN2	WSM6	KF	RRTMG	NU	662	16
11	070806040402	ACM2	Thompson	Tiedtke	RRTMG	NU	662	16
12	010602040402	YSU	WSM6	BMJ	RRTMG	NU	674	19
13	010601040402	YSU	WSM6	KF	RRTMG	NU	702	23
14	010606040402	YSU	WSM6	Tiedtke	RRTMG	NU	759	25
15	050606040402	MYNN2	WSM6	Tiedtke	RRTMG	NU	766	27
16	050602040402	MYNN2	WSM6	BMJ	RRTMG	NU	811	31
17	010806040402	YSU	Thompson	Tiedtke	RRTMG	NU	830	34
18	050806040402	MYNN2	Thompson	Tiedtke	RRTMG	NU	857	35

451 This gives two sets of physics combinations for additional testing: 1) one replaces only RRTMG  
 452 ((04|04)) for short and longwave radiation with New Goddard ((05|05)) making no other changes; and  
 453 2) RRTMG radiation is retained, but Noah-MP ((04)) replaces Noah-Unified ((02)). This creates an

454 additional 24 RCM configurations for assessment, bringing the total RCMs tested to 60. Although  
 455 Noah-MP has several parameter options, Phase II uses its default settings.

### 456 **3.4.3.3 Phase III (N=78): parameterising Noah-MP**

457 Phase II shows that RCM performance using New Goddard radiation is generally inferior to the same  
 458 RCMs using RRTMG (see [sSect. 5](#). RCM Physics test results). Consequently, RRTMG radiation is  
 459 re-adopted for Phase III. Conversely, a general performance improvement is conferred by using Noah-  
 460 MP over Noah-Unified ([sSect. 5](#)). Given this performance improvement using Noah-MP with default  
 461 settings, Phase III assesses RCM performances using specific parameter settings for Noah-MP.

462 Noah-MP provides a ‘dynamic vegetation cover’ model option (referred to as dynamic vege-  
 463 tation in the WRF users’ guide) (Niu et al., 2011). When deactivated (the default), monthly leaf area  
 464 index (LAI) is prescribed for various vegetation types and the greenness vegetation fraction (GVF)  
 465 comes from monthly GVF climatological values. Conversely, when dynamic vegetation cover is acti-  
 466 vated, LAI and GVF are calculated using a dynamic leaf model. We clarify here that dominant plant-  
 467 functional types do not change when using this option, but only the LAI and GVF, *i.e.*, only the  
 468 amount of green cover changes.

469 Noah-MP also provides several options for modelling surface run-off and groundwater pro-  
 470 cesses including a TOPMODEL (TOPography based hydrological MODEL)-based surface runoff  
 471 scheme and a simple groundwater model (SIMGM; Niu et al., 2011). Some studies have shown [that](#)  
 472 using this option improves [the](#) modelling of soil moisture (e.g. Zhuo et al., 2019). Thus, three new sets  
 473 of physics configurations are tested using Noah-MP where default options for specific settings are  
 474 changed as follows:

- 475 [3.4.](#) activate dynamic vegetation cover (dveg=2 in the WRF namelist); no other changes.
- 476 [4.5.](#) activate TOPMODEL runoff with simple groundwater (opt\_run=1); no other changes.
- 477 [5.6.](#) activate both dynamic vegetation and TOPMODEL runoff with simple groundwater; no other  
 478 changes.

479 As above, the worst performing RCMs in Phase II are excluded from Phase III testing. Based  
 480 on the RCM configuration performance rankings (Table 5), there is a sharp reduction in performance  
 481 starting from RCM #7 inclusive. Therefore, RCMs #7-12 are excluded from further testing. Phase III  
 482 thus comprises 18 new test simulations (sets 1-3 each comprising 6 RCMs) bringing the total RCMs  
 483 tested to N=78. Phase III physics tests are denoted using the same RCM identification schemes distin-  
 484 guished by appending ‘set\_1’, ‘set\_2’, ‘set\_3’ to identifiers.

485 **Table 5.** RCM physics combination ranks of the Phase II Noah-MP RCMs. Scores/ranks are based on model  
 486 bias and root mean square error for annual and seasonal precipitation, minimum temperature, maximum temper-  
 487 ature, climate extremes (wettest and hottest days), and Perkins Skill Scores (see [sSect. 3](#)).

No.	Physics combination	Rank total
-----	---------------------	------------

1	50801040404	721
2	70806040404	822
3	50802040404	848
4	70802040404	872
5	70601040404	880
6	50601040404	891
7	10802040404	988
8	70602040404	1005
9	70606040404	1028
10	10801040404	1042
11	70801040404	1056
12	10602040404	1264

---

488 **43.3.4 Shortlisting Physics Test RCMs for ERA5-~~NARCIIM2.0~~NARCIIM 2.0 evaluation**  
489 **simulations**

490 Considering the complete ~~NARCIIM2.0~~NARCIIM 2.0 N=78 physics test ensemble, to identify phys-  
491 ics test RCMs that perform poorly overall, RCMs are eliminated if they are in the lowest 1/3 for RCM  
492 performance ranks for any of maximum temperature, minimum temperature, precipitation, or for the  
493 overall model performance rank across these variables (see ~~s~~Sect. 5. RCM Physics test results). Under  
494 this scheme, 20 RCMs remain. The independence measures are then applied to the remaining 20  
495 RCMs to choose a final subset of 7 RCMs for ERA5-forced evaluation simulations (see ~~s~~Sect. 4.4).  
496 The ensemble size limit of N=7 is determined by available compute resources. These 7 candidate  
497 RCMs are assessed for potential use in the CMIP6 GCM-forced downscaling phase of ~~NAR-~~  
498 ~~CIIM2.0~~NARCIIM 2.0 (~~s~~Sect. 4.4 and Di Virgilio et al. in review).

499 **34.4 CORDEX ERA5-~~NARCIIM2.0~~NARCIIM 2.0 evaluation simulations**

500 NARCIIM1.x performed production climate simulations using a two-phase process. Its RCM physics  
501 testing selected definitive ‘production-grade’ RCMs which were then used to downscale both reanaly-  
502 sis data and CMIP3/5 GCMs. In contrast, for ~~NARCIIM2.0~~NARCIIM 2.0, as described above the  
503 N=78 RCM physics testing culminates in shortlisting 7 ‘production-candidate’ RCMs which are used  
504 to downscale the ERA5 reanalysis for 42-years (1979-2020). This enables assessment of ~~shortlisted~~  
505 ~~RCM-the~~ performances of these 7 shortlisted RCMs over a climatological period rather than the single  
506 year (2016) of the physics testing, which helps ascertain that performance differences between  
507 shortlisted RCMs are robust across a multi-decadal timescale capturing climatologically diverse years.  
508 The aim is that two definitive production-grade RCMs can be selected for CMIP6-forced downscaling  
509 from these ERA5-forced CORDEX ‘evaluation’ simulations. Thus, the seven ERA5-  
510 ~~NARCIIM2.0~~NARCIIM 2.0 RCMs were driven by ERA5.0 boundary conditions for January 1979 to

511 December 2020 using the model and nested domain setups described above for ~~NAR-~~  
512 ~~CM2.0~~NARCM 2.0. The skill of these RCMs in simulating the recent Australian climate was as-  
513 sessed as follows (see Di Virgilio et al. in review): annual and seasonal means were calculated for  
514 maximum and minimum temperature and precipitation using monthly means for temperature varia-  
515 bles, and the monthly sum for precipitation. Extremes of maximum temperature and precipitation (99<sup>th</sup>  
516 percentiles) and extreme minimum temperature (1<sup>st</sup> percentile) were calculated using daily data. RCM  
517 performances in reproducing observations over these timescales were assessed by calculating model  
518 outputs minus observations (~~i.e.~~, model bias), and the RMSE of modelled versus observed fields.  
519 RCM skill in simulating distributions of observed variables was assessed by comparing the PDFs for  
520 daily mean observations versus those of the RCMs. The ultimate outcome of these ERA5-forced sim-  
521 ulations and their evaluation is the selection of two definitive RCM configurations, R3 and R5, to run  
522 the CMIP6-forced phase of ~~NARCM2.0~~NARCM 2.0, see Di Virgilio et al. (in review) for further  
523 details on the evaluation methods and results. Supporting Information Figure S1 shows the WRF  
524 namelist settings for the R3 and R5 RCMs (see also sSect. 9. Code Availability).

### 525 **43.5 CORDEX CMIP6-forced ~~NARCM2.0~~NARCM 2.0 simulations**

526 The ideal CMIP6 GCM variables and their frequencies required to run the WRF RCM are listed in  
527 Table S1. A minority of variables in Table S1 are not available at sub-daily frequencies for every tar-  
528 get GCM. This necessitates assumptions/data proxies to be made. For instance, soil moisture and soil  
529 temperature variables were unavailable for some selected GCMs; hence, surrogate data, such as sur-  
530 face temperature, were used for initialisation (noting that soil data are only used by the RCM at ini-  
531 tialisation). In these cases, we investigated how long it took for uncertainty in the initial conditions to  
532 disappear from the WRF output by analysing the regionally averaged soil moisture time series. The  
533 data were regionalised according to the four Australian Natural Resource Management (NRM) re-  
534 gions / climate zones (Supporting Information Figure S2) which are broadly aligned with climatologi-  
535 cal boundaries (Fiddes et al., 2021) and with the IPCC reference regions (Iturbide et al., 2020). Time  
536 series plots (Figure S3) show that soil moisture equilibrates to be within a normal range following  
537 initialisation, indicating that the 12-month spin-up year (1950) is sufficient to account for the assump-  
538 tions made at model initialisation.

539 Boundary and initial conditions were prepared using selected GCM data to run the 151-year  
540 GCM-driven simulations using WRF version 4.1.2. The GCM-driven simulations were run and com-  
541 pleted using the pre-defined RCM settings for the two definitive RCM configurations using the WRF  
542 namelists in Supporting Information Figure S1 (see also sSect. 9. Code Availability). A cold restart  
543 was performed on the last Historical experiment year (2014), thus enabling the SSP1-2.6 and SSP3-  
544 7.0 experiments to be run for 2015-2100 concurrently with the Historical experiment. Testing the time

545 [duration required for soil moisture to equilibrate from the cold start showed that 1 year is sufficient.](#)

546 The 2014 cold start year is eventually overwritten by Historical runs initiated in 1950.

## 547 **6.5. RCM Physics test results**

### 548 **5.1 Phase I RCM performance summary**

549 The spatial variation and magnitudes for Phase I RCM biases and RMSEs for annual mean maximum  
 550 and minimum temperature and precipitation are shown in Figures 4-5, respectively. Overall, RCMs  
 551 are biased cold for maximum temperature (mean absolute bias for the ensemble mean = 1.18 K), and  
 552 warm-biased for minimum temperature (mean absolute bias = 1.31 K; Figure 4a-b). Maximum tem-  
 553 perature RMSE magnitudes are large over the elevated terrain of the southeast coast and over western  
 554 regions (Figure 5a). The simulation of precipitation shows biases of varying sign, with wet biases that  
 555 are strongest over eastern coastal regions (Figure 4c). Precipitation RMSEs are particularly large  
 556 along the eastern coastline (>15 mm), and generally show an east-west gradient, *i.e.*, progressively  
 557 decreasing further inland from the coast (Figure 5c).

### 558 **5.2 Comparing Phase II Physics Test RCM performances versus Phase I**

#### 559 **5.2.1 Climate Means**

560 Overall, the RCM ensemble using New Goddard (NG) radiation has inferior performance to the corre-  
 561 sponding RCMs using RRTMG in terms of annual/seasonal mean maximum temperature biases,  
 562 RMSEs, and PSS (Table 7). In contrast, NG confers superior performance for annual/seasonal mean  
 563 minimum temperature for these statistics. RCMs using NG show reduced biases for annual mean and  
 564 spring-time precipitation, but larger errors for DJF and JJA (Table 7). RMSEs for annual and seasonal  
 565 precipitation are similarly variable.

566 **Table 7.** Climate means performance: phase II physics tests (*i.e.*, N=12 set 1 changing only RRTMG to New  
 567 Goddard (NG) and N=12 set 2 changing only land surface model (LSM) from Noah-Unified to Noah-MP  
 568 (NMP) compared with the phase I physics test RCMs that were shortlisted for further testing (N=12).

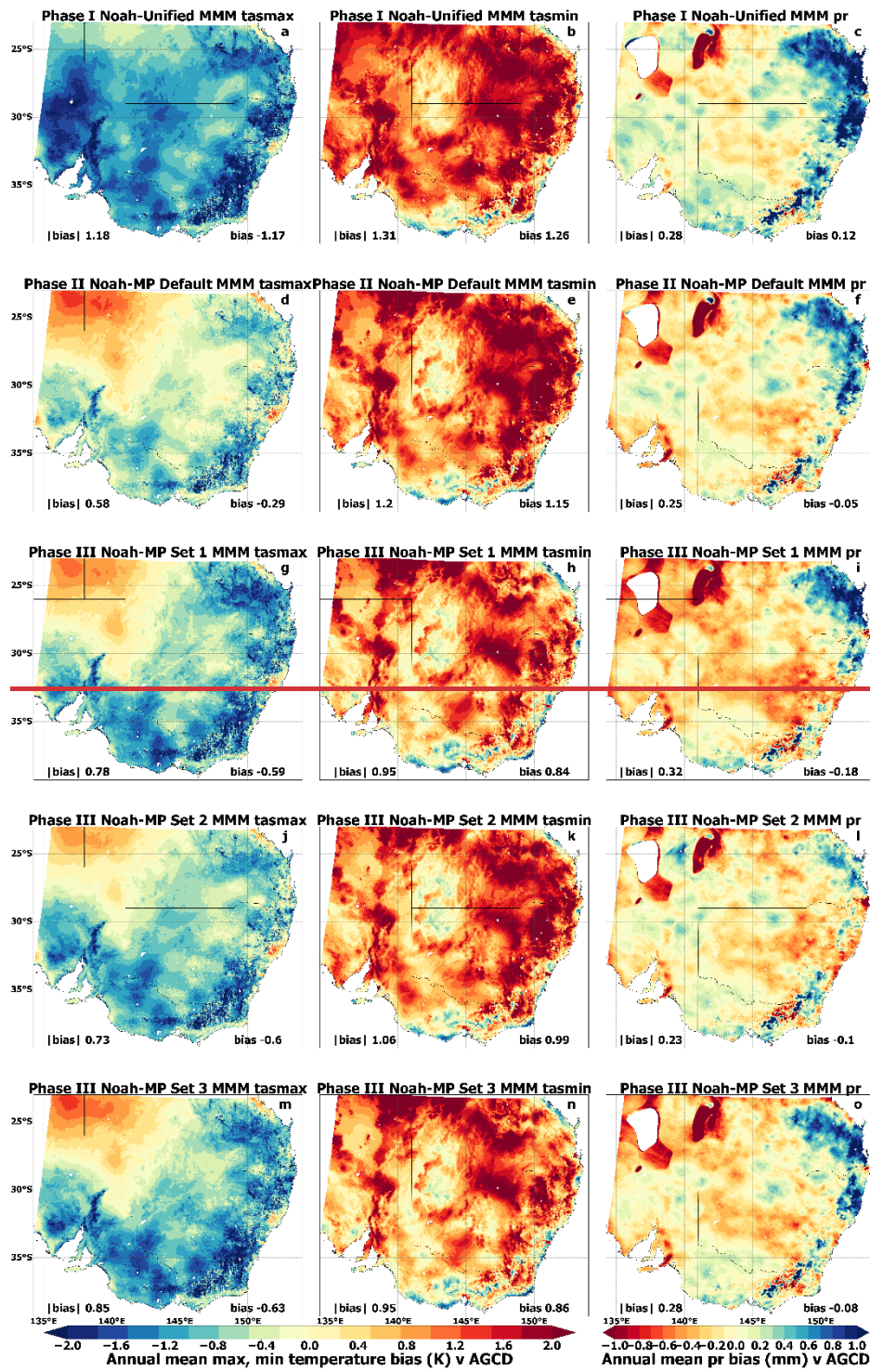
Variable	Timescale	Bias			RMSE			PSS		
		Phase I (N=12) ensemble mean	Phase II (NG rad.) ensemble mean	Phase II (NMP LSM) ensemble mean	Phase I (N=12) ensemble mean	Phase II (NG rad.) ensemble mean	Phase II (NMP LSM) ensemble mean	Phase I (N=12) ensemble mean	Phase II (NG rad.) ensemble mean	Phase II (NMP LSM) ensemble mean
Temp. Max. (K)	Annual	0.87	1.27	0.58	3.56	3.73	3.50	<b>0.950</b>	<b>0.936</b>	<b>0.955</b>
	DJF	0.74	1.29	0.63	4.41	4.70	4.43	-	-	-
	MAM	1.40	2.06	0.83	3.68	3.92	3.55	-	-	-

<b>Temp.</b> <b>Min. (K)</b>	JJA	0.62	0.81	0.52	2.64	2.66	2.65	<b>0.927</b>	<b>0.941</b>	<b>0.931</b>
	SON	0.87	1.04	0.66	3.25	3.32	3.20			
	Annual	1.35	0.95	1.2	3.53	3.41	3.42			
	DJF	1.50	1.08	0.87	3.86	3.82	3.66			
	MAM	1.21	0.84	0.92	3.55	3.45	3.50			
	JJA	0.82	0.51	0.91	3.00	2.92	3.00			
<b>Prec.</b> <b>(mm)</b>	SON	1.88	1.47	1.92	3.63	3.40	3.58	-	-	-
	Annual	0.25	0.24	0.25	7.21	7.32	6.78	<b>0.943</b>	<b>0.950</b>	<b>0.946</b>
	DJF	0.41	0.53	0.49	8.28	8.83	8.85			
	MAM	0.32	0.32	0.25	5.91	6.47	5.53			
	JJA	0.37	0.53	0.44	7.63	7.34	7.65	-	-	-
	SON	0.34	0.22	0.39	6.68	6.18	6.92			

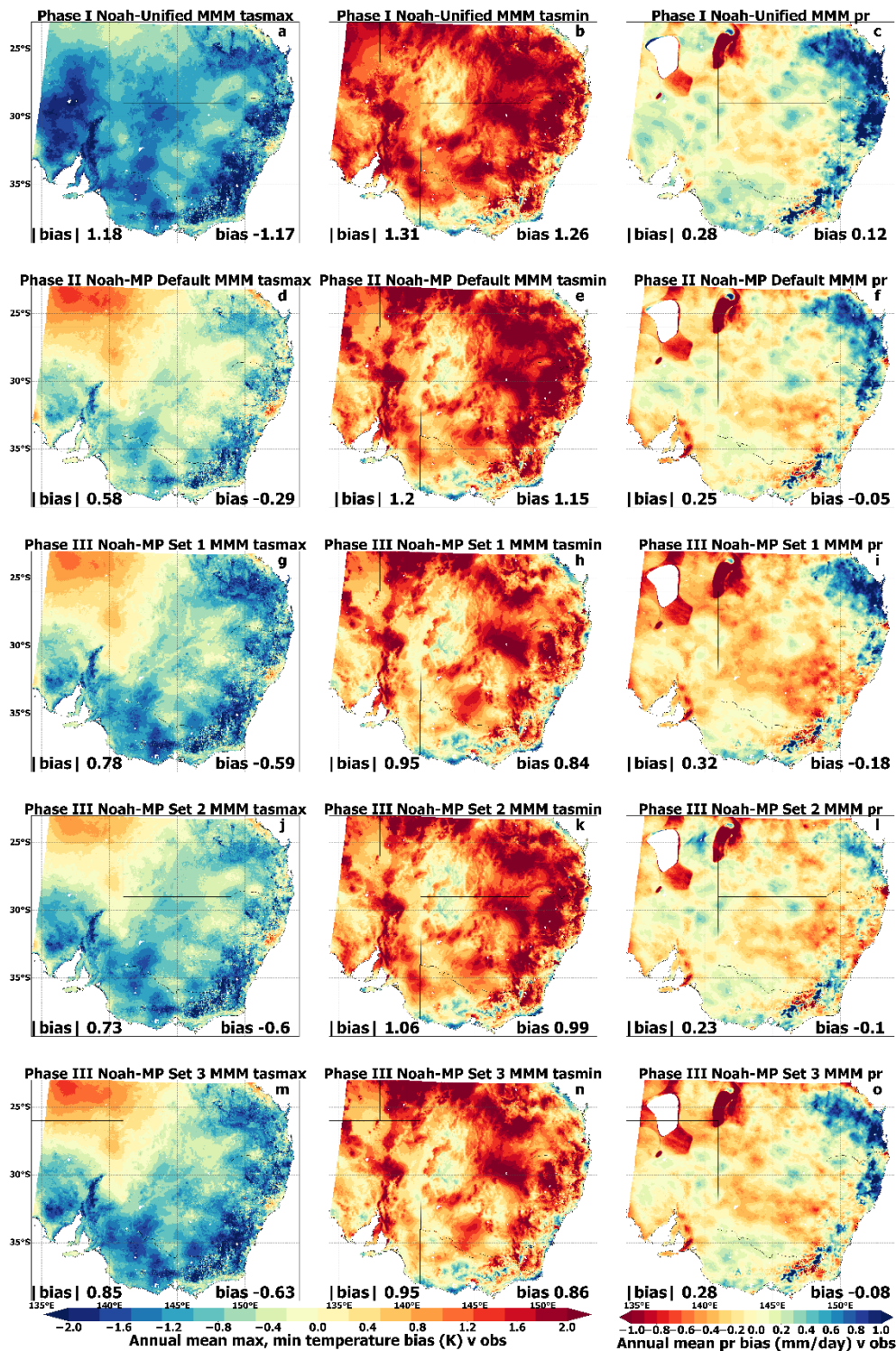
569 Phase II RCMs using Noah-MP with RRTMG retained show improved performance in simu-  
570 lating mean maximum and minimum temperature at annual timescales and most seasons relative to  
571 corresponding Phase I RCMs using Noah-Unified (Table 7; Figure 4-5). For instance, the mean abso-  
572 lute bias for annual mean maximum temperature is 0.58 K for the Noah-MP ensemble mean versus  
573 1.18 K for the Noah-Unified ensemble. In particular, cold bias magnitudes for maximum temperature  
574 are considerably lower over eastern and southern regions for the RCMs using Noah-MP (Figure 4d).  
575 RMSE magnitudes for maximum temperature are substantially reduced over the topographically com-  
576 plex regions of the southeast, and southwest and central regions (Figure 5d).

577 Overall, the magnitude of warm biases for minimum temperature are broadly similar for  
578 Phase I and Phase II RCMs (Figure 4b,c). Conversely, while RCMs in both Phases show large  
579 RMSEs for minimum temperature over several eastern regions, RMSEs are smaller for the Noah-MP  
580 ensemble over some southern areas (Figure 5b,c).

581 In contrast to the above results for the simulation of maximum temperature, overall, Phase II  
582 RCMs using Noah-MP show smaller performance improvements for the simulation of precipitation  
583 relative to the Phase I RCMs (Table 7). However, precipitation bias magnitudes are smaller for the  
584 Noah-MP ensemble over specific regions, e.g., north-eastern coastal regions and the elevated terrain  
585 of the south-east (Figure 4c,f).



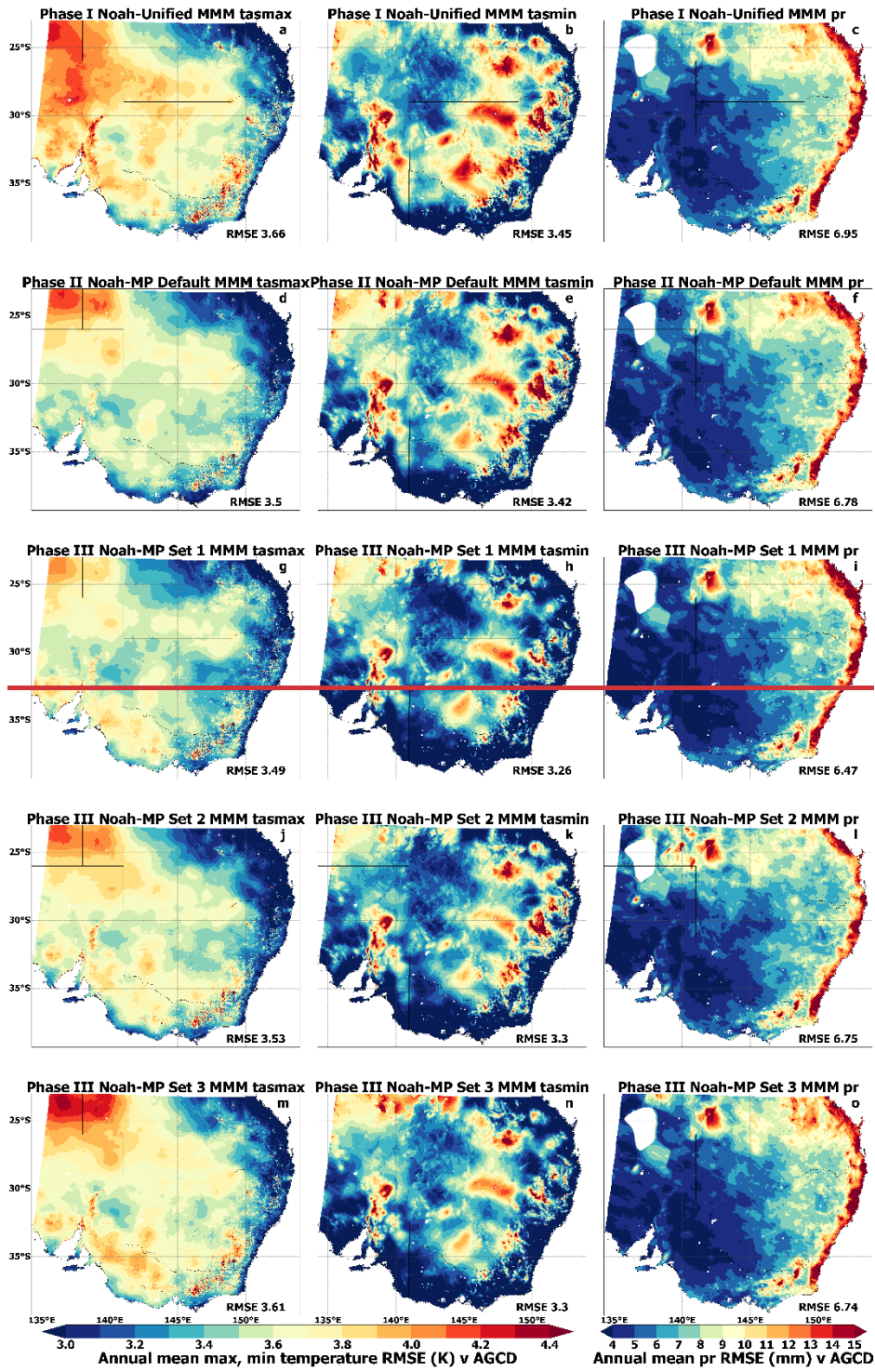


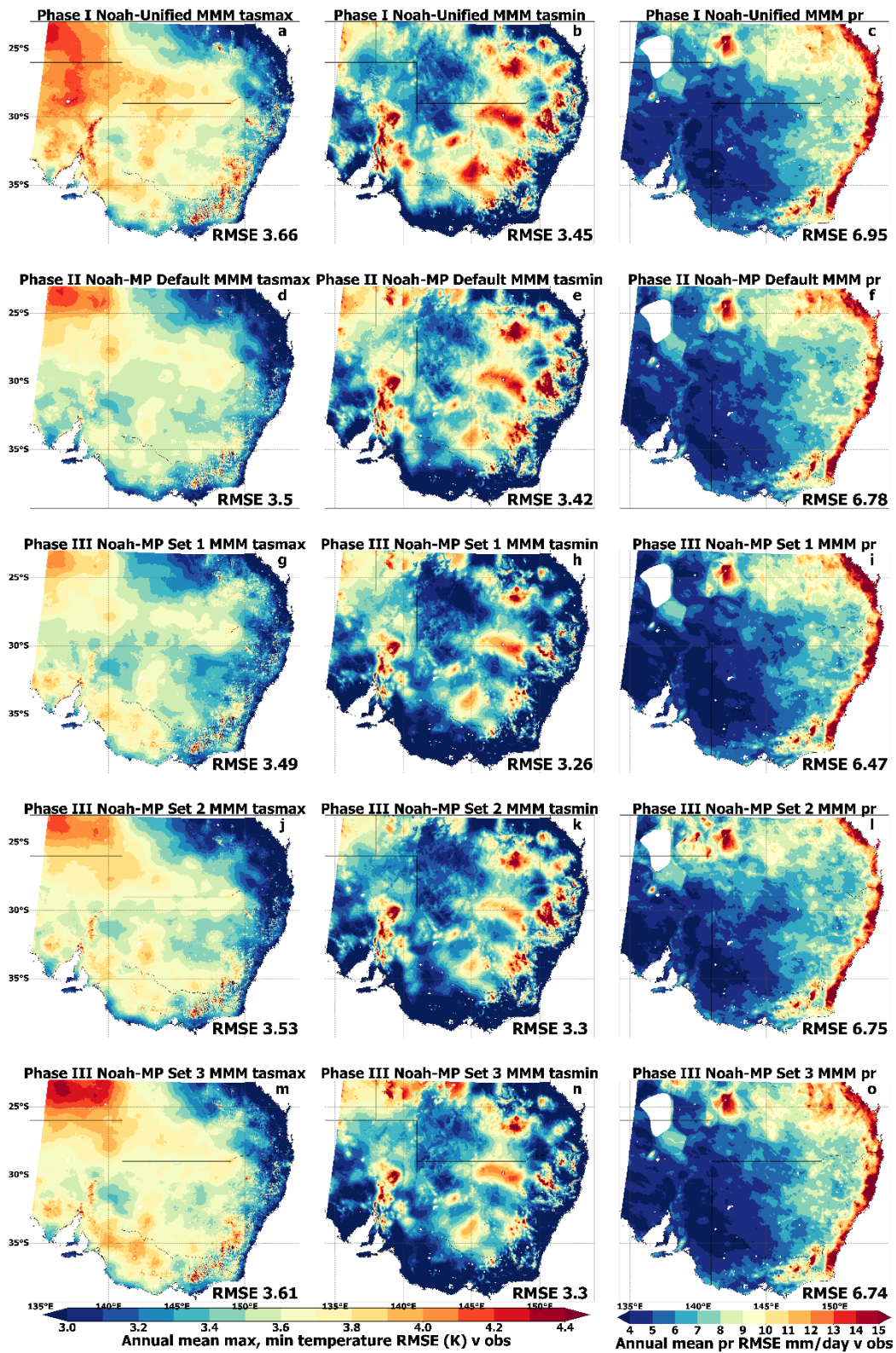


587

588 **Figure 4.** Phase I (N=36), Phase II (N=60) and Phase III (N=78) ensemble mean biases for annual mean maxi-  
 589 mum temperature, minimum temperature and precipitation with respect to Australian Gridded Climate Data  
 590 (AGCD) observations for ~~NARCIv2.0~~NARCIv2.0 Phase I physics test RCMs using Noah-Unified as the  
 591 land surface model (LSM) (a-c); Phase II physics test RCMs using Noah-MP as the LSM and its default settings  
 592 (d-f); Phase III <sup>1</sup>set physics test RCMs using Noah-MP with dynamic vegetation cover activated (g-i); Phase  
 593 III <sup>2</sup>set physics test RCMs using Noah-MP with TOPMODEL surface runoff and simple groundwater activat-

594 ed (j-l); and Phase III ~~set 3~~ physics test RCMs using Noah-MP with both dynamic vegetation cover and TOP-  
595 MODEL runoff activated (m-o).





597

598 Figure 5. As per Figure 4 but showing RMSEs.

599 **5.2.2. Climate Extremes**

600 Climate extreme analysis assesses RCM representations of the hottest and the wettest day versus

601 AGCD. For both extremes and for RCM biases and RMSEs, Phase II RCMs using NG radiation

602 showed inferior performance relative to phase I RCMs using RRTMG (Table 8). Conversely, Phase II  
 603 RCMs using Noah-MP show substantial reductions in bias for both the hottest and wettest days (Table  
 604 8). Phase II Noah-MP RCMs show a small increase in RMSE for the hottest day (Phase I bias=3.59  
 605 K; Phase II bias=3.74 K); however, RMSEs are smaller for the wettest day (~~i.e.~~ Phase I  
 606 RMSE=19.20 mm; Phase II RMSE=18.47 mm) (Table 8).

607 **Table 8** Climate extremes performance: comparing phase I RCMs (N=12) with phase II RCMs  
 608 (~~i.e.~~ 12 RCMs changing radiation from RRTMG to New Goddard (NG) and 12 RCMs changing  
 609 land surface model (LSM) from Noah-Unified to Noah-MP; NMP).

Variable	Bias			RMSE		
	Phase I (N=12) ensemble mean	Phase II (NG rad.) ensemble mean	Phase II (NMP LSM) ensemble mean	Phase I (N=12) ensemble mean	Phase II (NG rad.) ensemble mean	Phase II (NMP LSM) ensemble mean
Temp. max: hottest (K)	1.11	1.93	0.81	3.59	3.97	3.74
Prec.: wettest (mm)	3.08	3.21	2.60	19.20	20.52	18.47

### 610 5.3 Phase III RCM performance summary and shortlisting N=7 RCMs for

#### 611 ERA5-~~NARCI~~NARCIM 2.0 evaluation simulations

612 Overall, RCM biases for mean maximum temperature do not show marked improvements once the  
 613 dynamic vegetation cover and surface runoff options are activated for Noah-MP (Figure 4 g,j,m) rela-  
 614 tive to RCMs using Noah-MP with default settings (Figure 4d). However, specifically for the RCM  
 615 ensemble with dynamic vegetation cover activated for Noah-MP, RMSE magnitudes for maximum  
 616 temperature are lower over some eastern coastal regions (Figure 5g).

617 The simulation of mean minimum temperature shows clear performance improvements for  
 618 Phase III RCMs using options activated for Noah-MP, relative to RCMs using Noah-MP defaults.  
 619 Overall, both biases and RMSEs for minimum temperature are reduced in magnitude for RCMs using  
 620 ~~the~~ either or both of dynamic vegetation cover and runoff/groundwater options activated for Noah-  
 621 MP, relative to the default parameters (Figure 4-5). These performance improvements are largest over  
 622 eastern and southern regions.

623 There are no substantial overall performance improvements in the simulation of precipitation  
 624 for Phase III RCMs relative to Phase II RCMs (Figures 4-5 f,i,l,o). However, using Noah-MP with

625 specific LSM options remains favourable to using RCMs with Noah-Unified, albeit the performance  
626 gains are generally small, except for some coastal regions and especially the north-east.

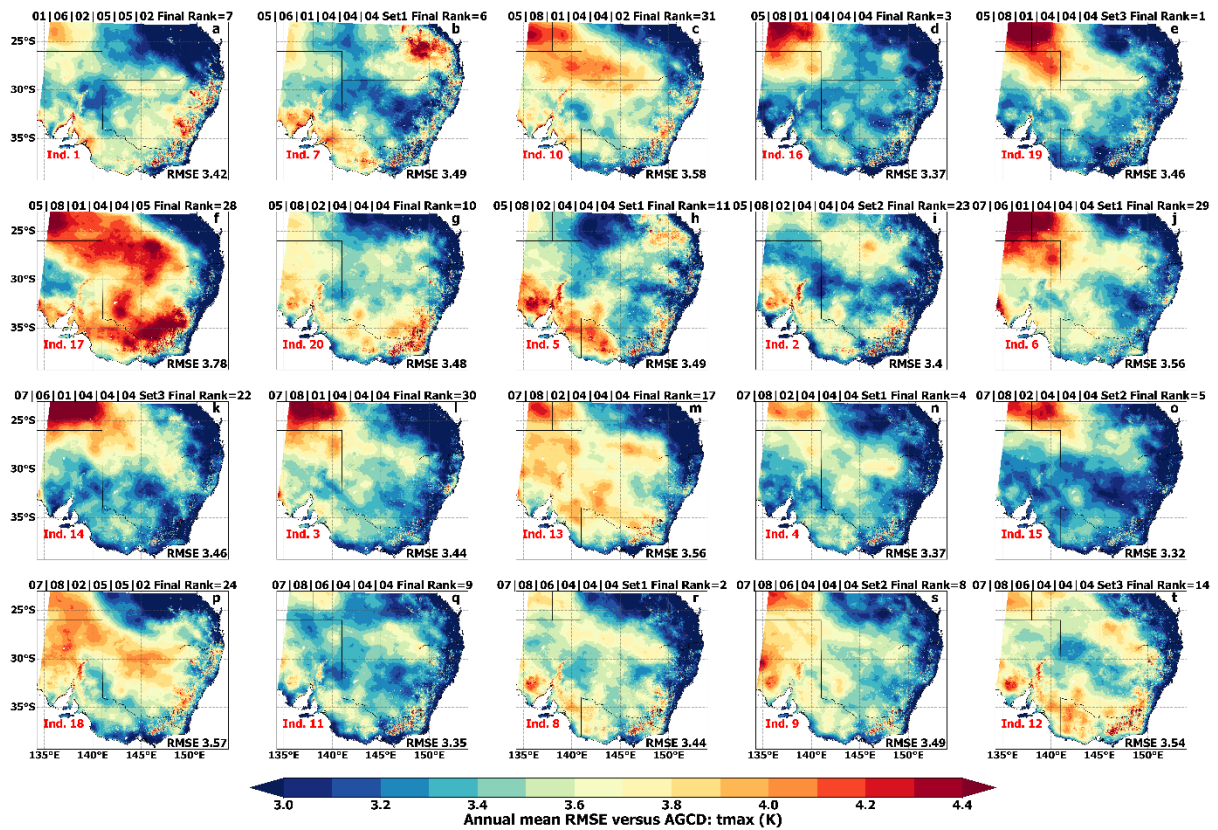
627 All 78 RCMs in the complete RCM physics test ensemble are ranked for performance as de-  
628 scribed in [Sect. 3.4.2](#). Once the poor-performing RCMs are excluded, there are 20 RCMs remaining  
629 (Table 9; Figures 6-8). In Table 9, we see that 16 Noah-MP-based RCMs from Phase II and Phase III  
630 comprise this set of 20 RCMs, with 3 of the 20 RCMs using Noah-Unified, and 1 using CLM4.0. For  
631 maximum temperature, some shortlisted RCMs show [large-substantial](#) RMSEs over north-western and  
632 inland areas (e.g., Figure 6 d-f) that are of [similar-larger](#) magnitude [over these areas to those of](#) than  
633 the ensemble means of Phase I-III RCMs (Figure 5). Conversely, several shortlisted RCMs show very  
634 low RMSEs for maximum temperature across eastern and southern regions, especially along the east-  
635 ern coast (Figure 6, e.g., RCMs in panels d,l,n,o,q). For minimum temperature, a subset of the twenty  
636 shortlisted RCMs show substantially reduced RMSEs over many regions relative to the Phase I-III  
637 ensemble means (Figure 7, e.g., RCMs in panels: b,h,i). Additionally, several shortlisted RCMs show  
638 reduced RMSEs for precipitation over the eastern coast and north-east (Figure 8, e.g., RCMs in pan-  
639 els: c, l, m, n, o) relative to the Phase I-III RCM ensemble means in Figure 5c,f,i,l,o.

640 These 20 RCMs are assessed for statistical independence and 7 RCMs from this RCM set are  
641 shortlisted for the ERA5-forced RCM simulations considering both their performance and independ-  
642 ence scores (Table 9). These 7 shortlisted RCMs are listed in **bold** in Table 9 and are identified as R1-  
643 R7 in the ERA5-forced evaluation simulations (Table 9; final column). RCMs are shortlisted from the  
644 set of 20 if they rank highly for both performance and independence. For instance, RCM  
645 050801040404\_set\_3 (top row, Table 9) is top-ranked for performance, however, its independence  
646 scores/ranks are low, hence it is not shortlisted. It is important to note that, while a general perfor-  
647 mance gain is observed in the physics testing when using Noah-MP, there are some specific RCM  
648 configurations using Noah-Unified that perform well in simulating the Australian climate. For in-  
649 stance, the RCM 010602050502 (row 7; Table 9; **R1**) uses Noah-Unified and performs well overall  
650 (its overall performance rank=7), and especially for the simulation of maximum temperature (Figure  
651 6a). It is also the only RCM in this set of 20 RCMs to use YSU for PBL. Importantly, this RCM is  
652 highly ranked for statistical independence, hence, this RCM is shortlisted for the N=7 set. [We note](#)  
653 [here that R1-R7 are simply a chronological naming convention and do not imply any ranking for these](#)  
654 [7 RCM configurations.](#)

655 **Table 9.** The 20 ~~NARCIIM2.0~~[NARCIIM 2.0](#) physics test RCMs shortlisted from the ensemble of 78 RCMs  
656 based on their performance in simulating the Australian climate and independence (Ind.). N=7 **R1-R7** RCMs  
657 shortlisted for ERA5-forced evaluation simulations shown in **bold**. **R1-R7** are [a naming convention and do not](#)

658 [imply a ranking for these 7 RCMs](#). NU=Noah Unified; NMP=Noah-MP; DV=dynamic vegetation cover;  
 659 TOP=topmodel runoff.

#	RCM Physics Combination	PBL	MP	Cumulus	SW/LW	LSM	Test Phase	Overall Performance Rank	Bishop Abramowitz Ind. Rank	Herger Ind. Set 1	Herger Ind. Set 2	ERA5-forced RCM Identifier
1	050801040404_set_3	MYNN2	Thom	KF	RRTMG	NMP DV+TOP	III	1	19	20	20	
2	<b>070806040404_set_1</b>	<b>ACM2</b>	<b>Thom</b>	<b>Td</b>	<b>RRTMG</b>	<b>NMP DV</b>	<b>III</b>	<b>2</b>	<b>8</b>	<b>5</b>	<b>6</b>	<b>R6</b>
3	50801040404	MYNN2	Thom	KF	RRTMG	NMP	II	3	16	12	13	
4	<b>070802040404_set_1</b>	<b>ACM2</b>	<b>Thom</b>	<b>BMJ</b>	<b>RRTMG</b>	<b>NMP DV</b>	<b>III</b>	<b>4</b>	<b>4</b>	<b>3</b>	<b>3</b>	<b>R5</b>
5	070802040404_set_2	ACM2	Thom	BMJ	RRTMG	NMP TOP	III	5	15	13	12	
6	<b>050601040404_set_1</b>	<b>MYNN2</b>	<b>WSM6</b>	<b>KF</b>	<b>RRTMG</b>	<b>NMP DV</b>	<b>III</b>	<b>6</b>	<b>7</b>	<b>10</b>	<b>10</b>	<b>R2</b>
7	<b>10602050502</b>	<b>YSU</b>	<b>WSM6</b>	<b>BMJ</b>	<b>NG</b>	<b>NU</b>	<b>II</b>	<b>7</b>	<b>1</b>	<b>3</b>	<b>3</b>	<b>R1</b>
8	<b>070806040404_set_2</b>	<b>ACM2</b>	<b>Thom</b>	<b>Td</b>	<b>RRTMG</b>	<b>NMP TOP</b>	<b>III</b>	<b>8</b>	<b>9</b>	<b>9</b>	<b>5</b>	<b>R7</b>
9	70806040404	ACM2	Thom	Td	RRTMG	NMP	II	9	11	14	14	
#	50802040404	MYNN2	Thom	BMJ	RRTMG	NMP	II	10	20	19	19	
#	<b>050802040404_set_1</b>	<b>MYNN2</b>	<b>Thom</b>	<b>BMJ</b>	<b>RRTMG</b>	<b>NMP DV</b>	<b>III</b>	<b>11</b>	<b>5</b>	<b>2</b>	<b>2</b>	<b>R3</b>
#	070806040404_set_3	ACM2	Thom	Td	RRTMG	NMP DV+TOP	III	14	12	10	10	
#	70802040404	ACM2	Thom	BMJ	RRTMG	NMP	II	17	13	15	15	
#	070601040404_set_3	ACM2	WSM6	KF	RRTMG	NMP DV+TOP	III	22	14	16	16	
#	<b>050802040404_set_2</b>	<b>MYNN2</b>	<b>Thom</b>	<b>BMJ</b>	<b>RRTMG</b>	<b>NMP TOP</b>	<b>III</b>	<b>23</b>	<b>2</b>	<b>4</b>	<b>4</b>	<b>R4</b>
#	70802050502	ACM2	Thom	BMJ	NG	NU	II	24	18	18	18	
#	50801040405	MYNN2	Thom	KF	RRTMG	CLM4	I	28	17	17	17	
#	070601040404_set_1	ACM2	WSM6	KF	RRTMG	NMP DV	III	29	6	7	8	
#	70801040404	ACM2	Thom	KF	RRTMG	NMP	II	30	3	1	1	
#	50801040402	MYNN2	Thom	KF	RRTMG	NU	I	31	10	6	7	



660

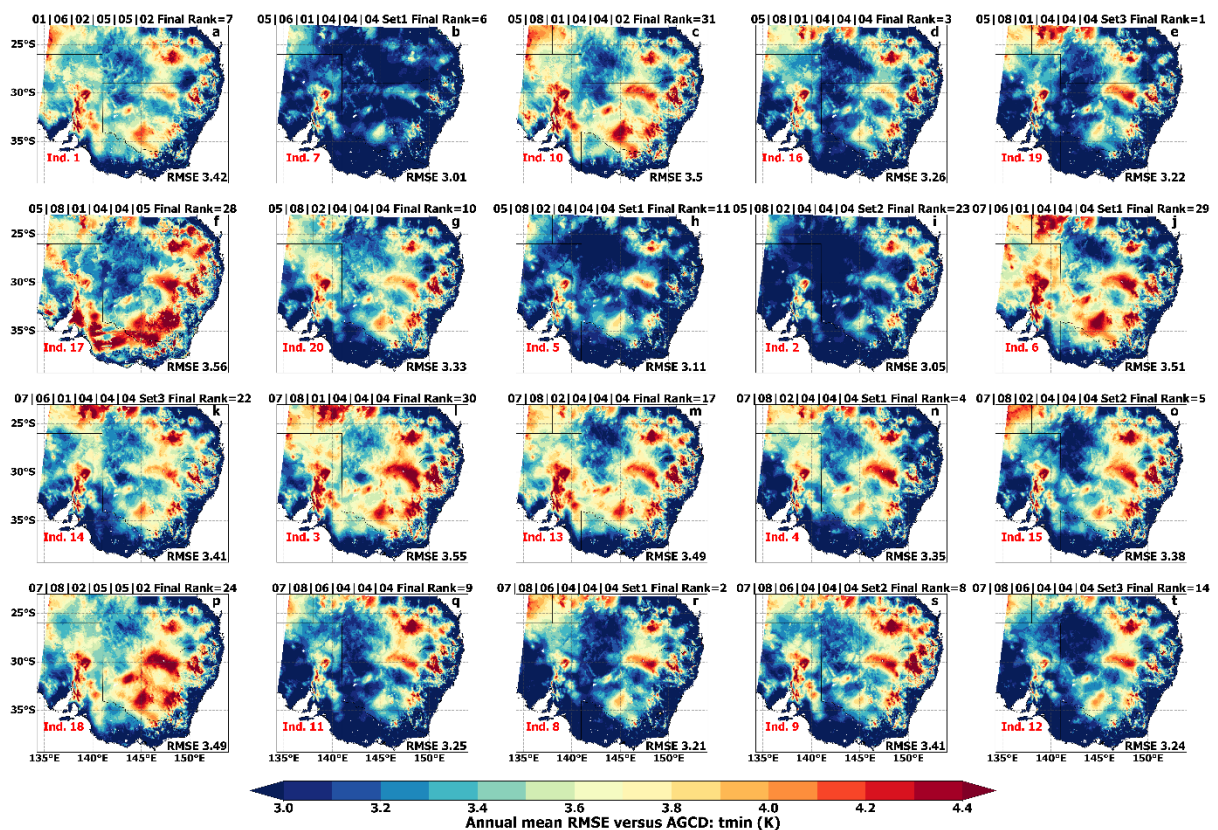
661

662

663

664

**Figure 6.** RMSEs for modelled mean maximum temperature (tmax) versus observations for the twenty [NARCIIM2.0](#) physics test RCMs shortlisted from the full ensemble of seventy-eight RCMs based on their performance in simulating the recent south-east Australian climate. Overall (final) performance ranks and Bishop and Abramowitz (2013) method independence (Ind.) scores are shown.

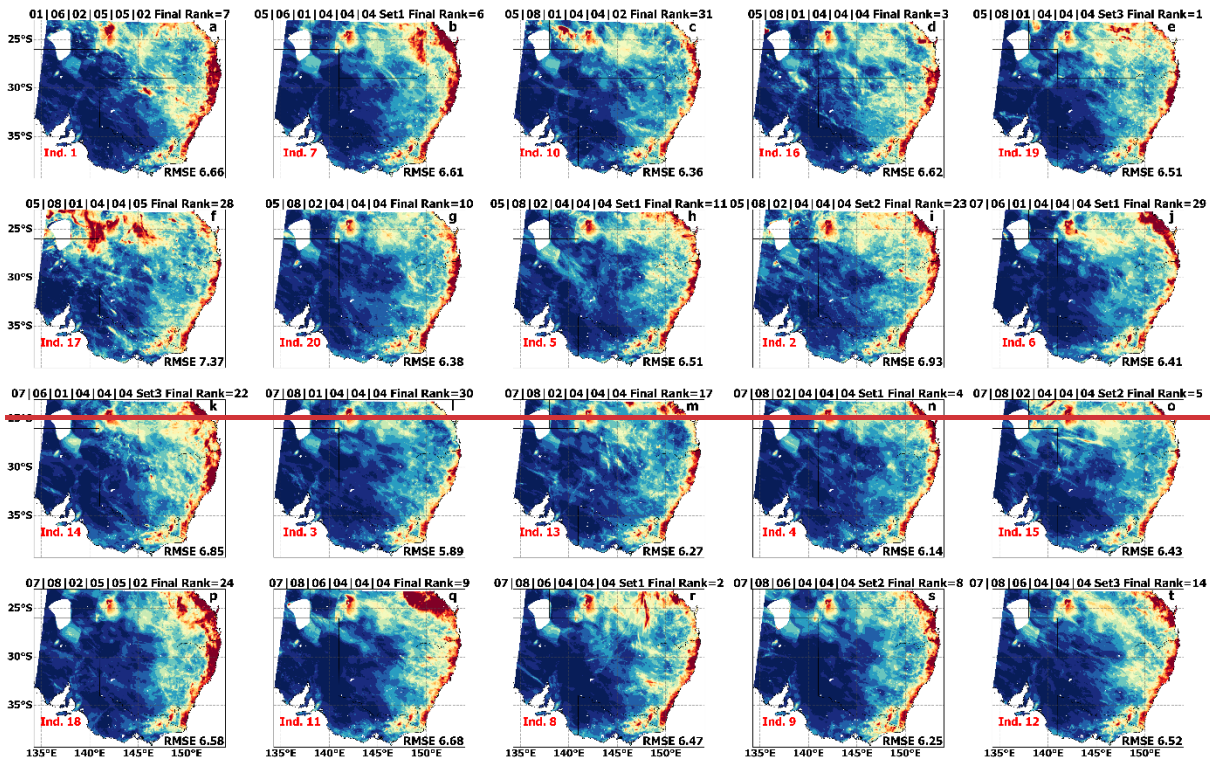


665

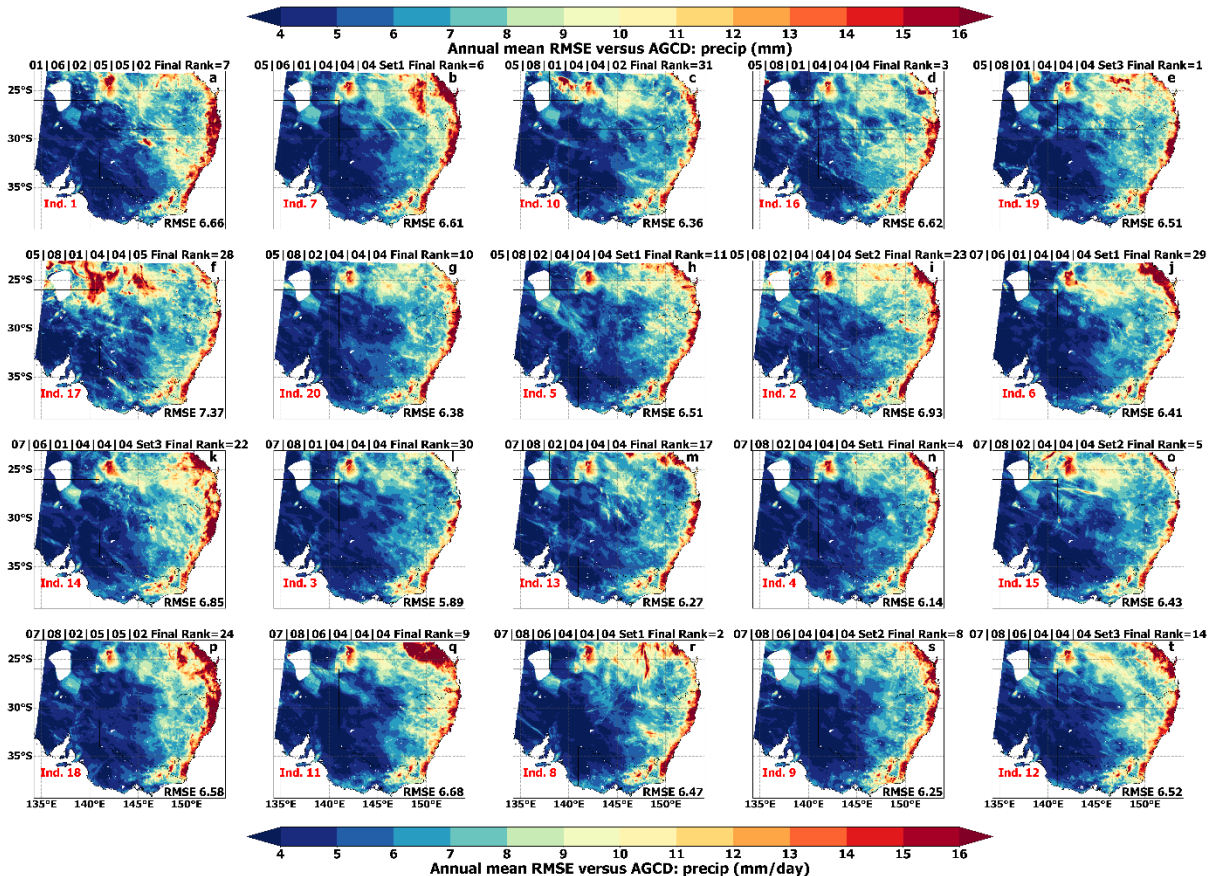
666

**Figure 7.** As per Figure 6 but for mean minimum temperature (tmin).





667



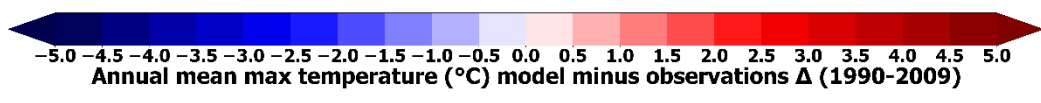
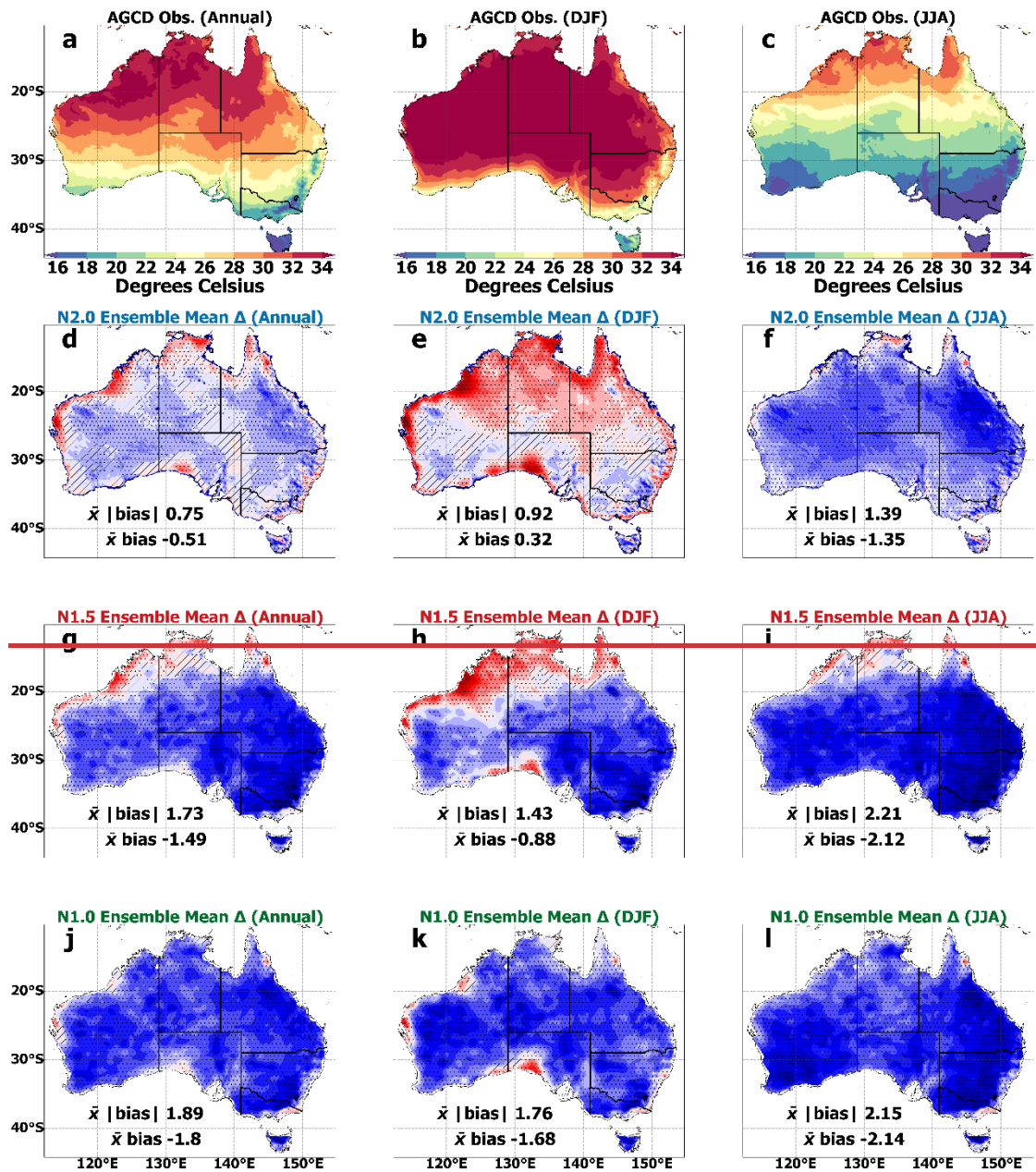
668

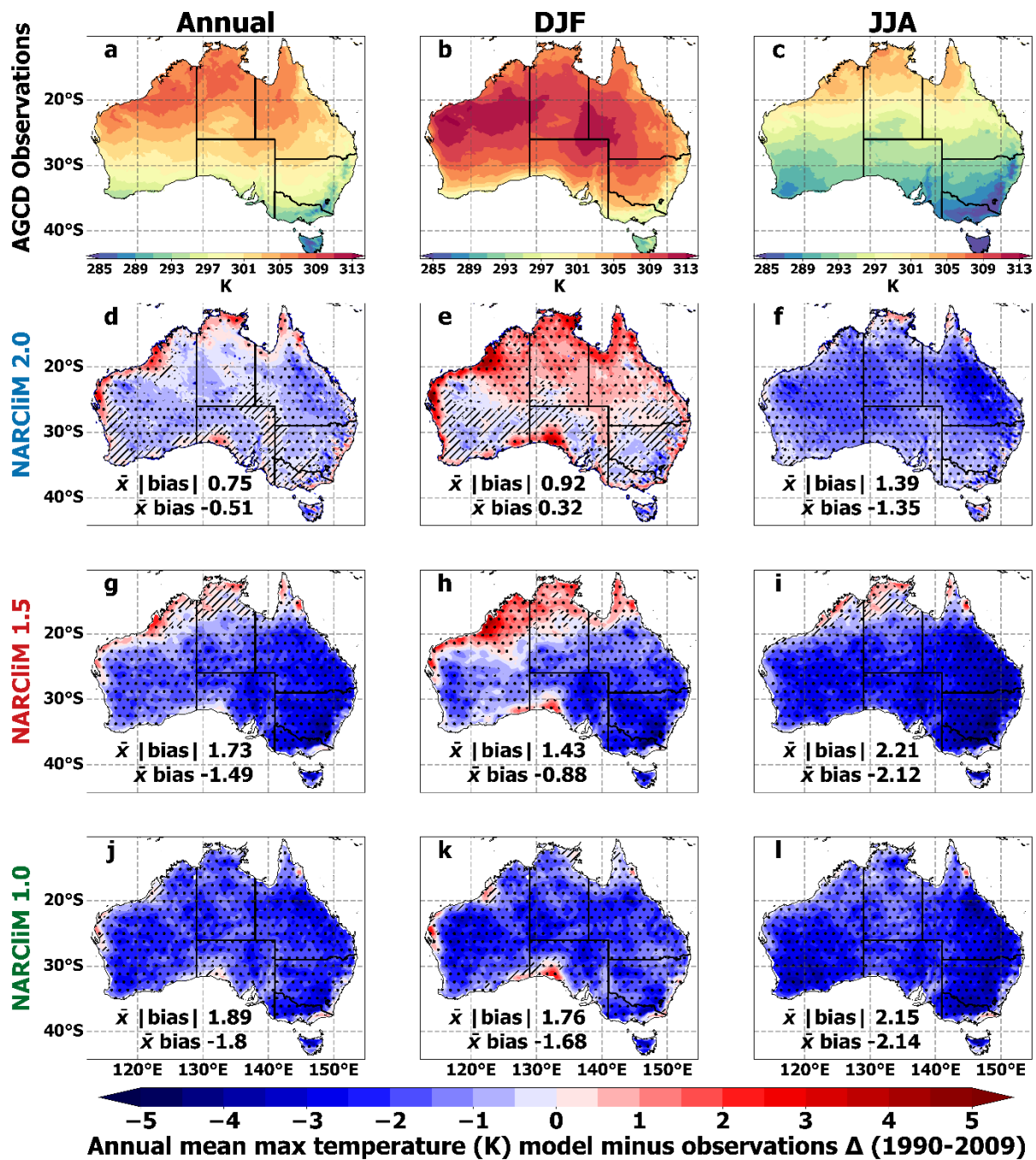
669 Figure 8. As per Figure 6 but for mean precipitation (precip.).

## 670 **7.6. CORDEX-CMIP6 ~~NARCIIM2.0~~NARCIIM 2.0 historical** 671 **evaluation**

### 672 **6.1 Maximum temperature**

673 ~~NARCIIM2.0~~Overall, NARCIIM 2.0 RCMs simulate maximum temperature more accurately than  
674 NARCIIM1.x, with widespread, statistically significant reductions in cold biases in the ensemble  
675 mean (Figure 9), as well as for many individual RCMs (Supporting Information Figure S4-S6). These  
676 reductions in bias apply for all timescales but are largest for the annual mean, i.e., the area-  
677 averaged mean absolute bias for the NARCIIM 2.0 ensemble is 0.75 K°C (range: 0.61 to 2.03 K)~~for~~  
678 ~~the NARCIIM2.0 ensemble~~, 1.73 K°C (range: 1.1 to 2.37 K) for ~~NARCIIM1.5~~NARCIIM 1.5, and  
679 1.89 °C K (range: 0.55 to 4.12 K) for ~~NARCIIM1.0~~NARCIIM 1.0 (Figure 9d,g,j and Figure S4). Nota-  
680 bly, the NARCIIM2.0 ensemble mean annual mean maximum temperature bias magnitudes are ~~very~~  
681 small, i.e., around <0.5 K°C, over south-west WA, southern coastal regions, and several eastern  
682 regions. This may be important from a climate change adaptation and mitigation perspective as these  
683 regions are heavily populated and economically significant. ~~NARCIIM2.0~~NARCIIM 2.0 retains warm  
684 biases of similar magnitude to ~~NARCIIM1.5~~NARCIIM 1.5 along the north-west coast of Australia  
685 (Figure 9d,g). Moreover, these warm biases cover additional areas for ~~NARCIIM2.0~~NARCIIM 2.0,  
686 especially during DJF (Figure 9e,h). ~~A~~ wide range of bias signs are evident for the individual ~~NAR-~~  
687 ~~CIIM2.0~~NARCIIM 2.0 ensemble members (Figures S4-S6) and a minority of NARCIIM 2.0 RCMs  
688 retain strong cold biases, e.g., at an annual timescale NARCIIM 2.0-NorESM2-MM R3 (mean abso-  
689 lute bias = 2.03 K) and UKESM-1-0-LL R3 (1.77 K). Additionally, ~~T~~the R5 RCM is generally warm-  
690 er than R3; (e.g., (Figure S4c,d). Considering the forcing GCM data, overall, ensemble means for the  
691 CMIP6 and CMIP5 GCMs generally show similar patterns and magnitudes of cold bias for maximum  
692 temperature (Supporting Information S7).





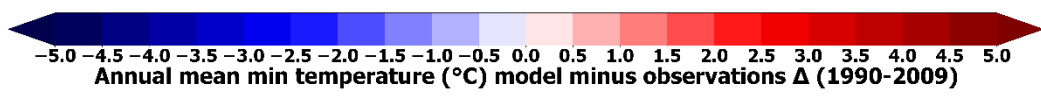
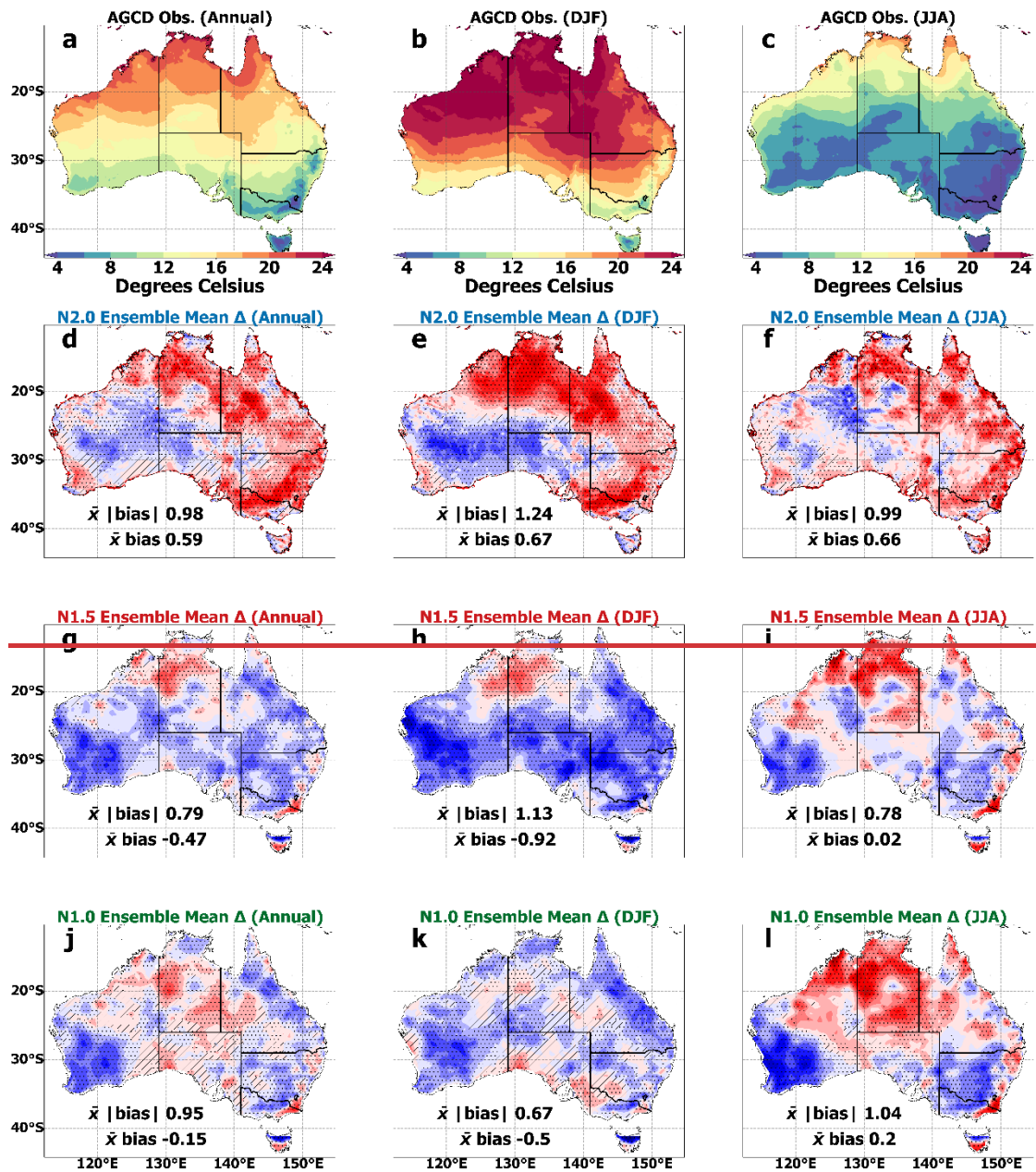
694

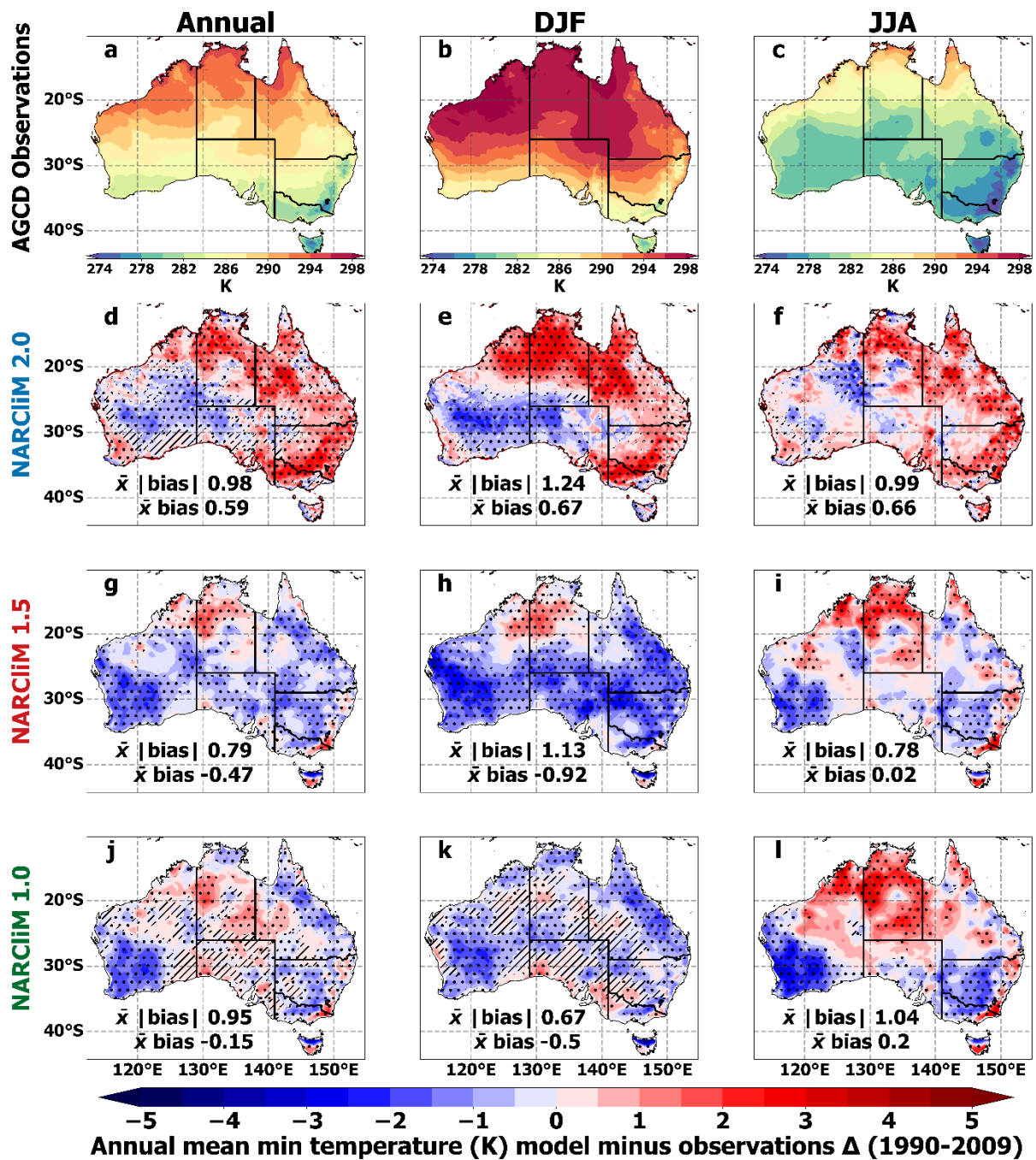
695 **Figure 9.** Annual, DJF and JJA mean near-surface atmospheric maximum temperature biases for ~~NAR-~~  
696 ~~CIIM 2.0~~ **NARCIIM 2.0**, 1.5 and 1.0 historical ensemble means with respect to Australian Gridded Climate Data  
697 (AGCD) observations for 1990-2009. Stippled areas indicate locations where an RCM shows statistically signif-  
698 icant bias. Significance stippling for the ensemble mean bias follows Tebaldi et al. (2011) and is applied sepa-  
699 rately to each RCM ensemble. Statistically insignificant areas are shown in colour, denoting that less than half  
700 of the models are significantly biased. In significant agreeing areas (stippled), at least half of RCMs are signifi-  
701 cantly biased, and at least 70% of significant RCMs in each ensemble agree on the direction of the bias. Signifi-  
702 cant disagreeing areas are shown in hatching, which are where at least half of the models are significantly biased  
703 and less than 70% of significant models in each ensemble agree on the bias direction - see main text for addi-  
704 tional detail on the stippling regime.

## 705 6.2 Minimum temperature

706 The simulation of mean minimum temperature by [NARCM2.0](#) is generally warm bi-  
707 ased at all timescales (Figure 10). Its bias magnitudes over many regions are larger versus [NAR-](#)  
708 [CM1.5](#), e.g., annual mean area-averaged absolute biases are 0.98 °C and 0.79 °C  
709 for [NARCM2.0](#) and [NARCM1.5](#), respectively (Figure 10 d,g). How-  
710 ever, there are exceptions to this result over specific regions, for example, parts of south-west western  
711 Australia show annual mean bias magnitudes of <1 °C for [NARCM2.0](#), but these  
712 areas show biases below -2 °C for [NARCM1.x](#) (Figure 10d,g,j). Most individual RCMs compris-  
713 ing the [NARCM2.0](#) ensemble show stronger warm biases than their [NAR-](#)  
714 [CM1.5](#) peers at both annual and seasonal timescales (Figures S8-S10). The ACCESS-  
715 ESM-1-5-forced [NARCM2.0](#) RCMs are considerably more warm-biased than the oth-  
716 er [NARCM2.0](#) RCMs, with average absolute biases of 1.74 °C and 1.9 °C; Fig.  
717 S8c-d).

718 Many of the CMIP6 GCMs used to force the [NARCM2.0](#) RCMs are warmer than  
719 the CMIP5 GCMs used to force [NARCM1.5](#), such that the ensemble mean bias of the  
720 former is 1.9 °C versus 1.11 °C (Figure S11). In particular, ACCESS-ESM-1-5 and MPI-ESM1-2-  
721 HR are substantially more warm-biased relative to all other selected GCMs, with mean absolute bias-  
722 es of 2.2 °C and 3.47 °C, respectively (Figure S11). This suggests that [NARCM2.0](#)'s warm biases for mean minimum temperature are at least partially inherited from the driving da-  
723 ta. However, whilst the ACCESS-ESM-1-5-forced [NARCM2.0](#) RCMs are much  
724 warmer than their counterparts (*i.e.*, 1.74 °C and 1.9 °C), this does not apply to the MPI-ESM1-  
725 2-HR-forced RCMs, which have biases of only 1.01 °C and 1.09 °C. Hence, factors additional to  
726 the driving data, such as changes in RCM parameterisations between [NARCM](#) generations and other  
727 model design changes likely contribute to the warmer biases observed for [NARCM2.0](#).  
728  
729





731  
732 Figure 10. As per Figure 9 but for mean minimum temperature.

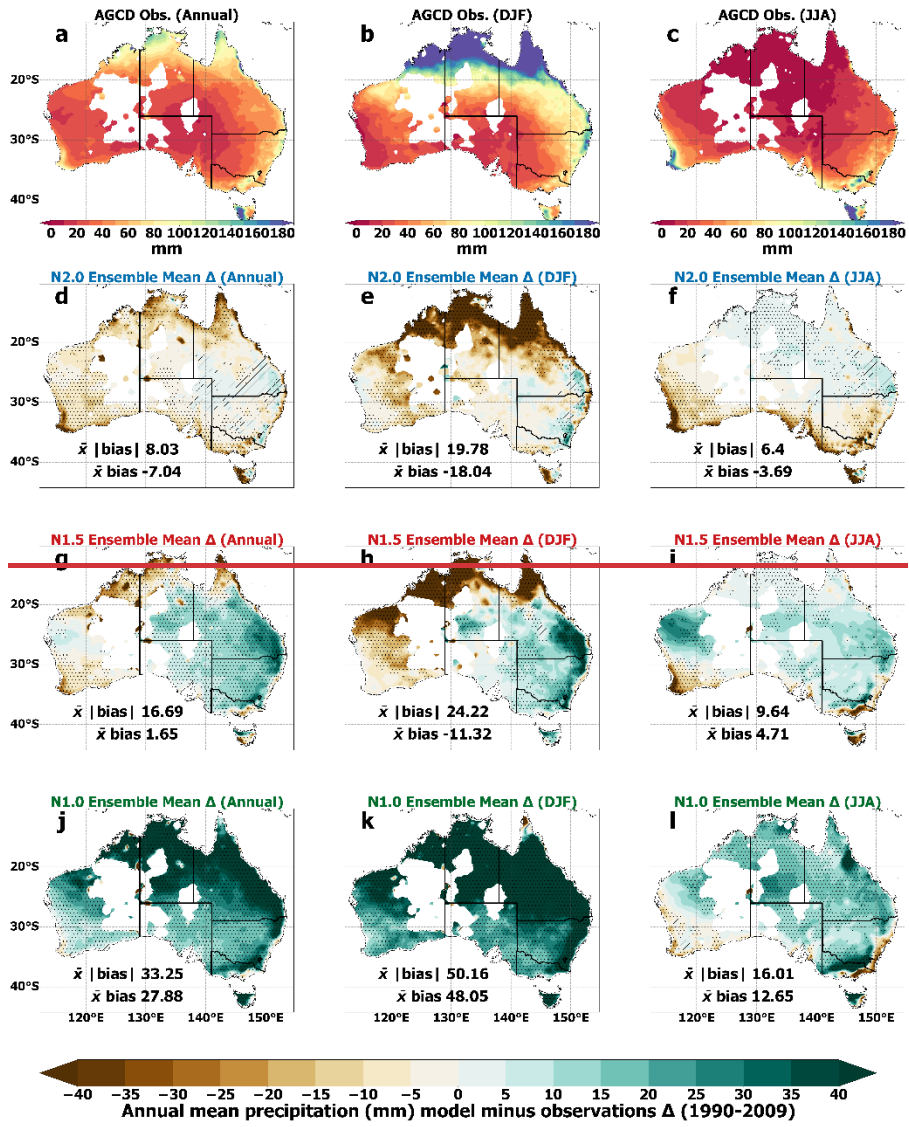
### 733 6.3 Precipitation

734 The [NARCIIM2.0](#) ensemble shows small dry biases for mean precipitation over most  
735 regions, except for some areas mainly in the east of the country which show slight wet biases (Figure  
736 11d-f). This contrasts with stronger, *statistically significant* wet biases of [NARCIIM1.5](#)  
737 *that are statistically significant over many regions* (Figure 11g-i) and the even stronger wet biases of  
738 [NARCIIM1.0](#) (Figure 11j-l). Area-averaged bias magnitudes are considerably smaller  
739 for [NARCIIM2.0](#) relative to NARCIIM1.x, especially for the annual mean, *i.e.* 8.03

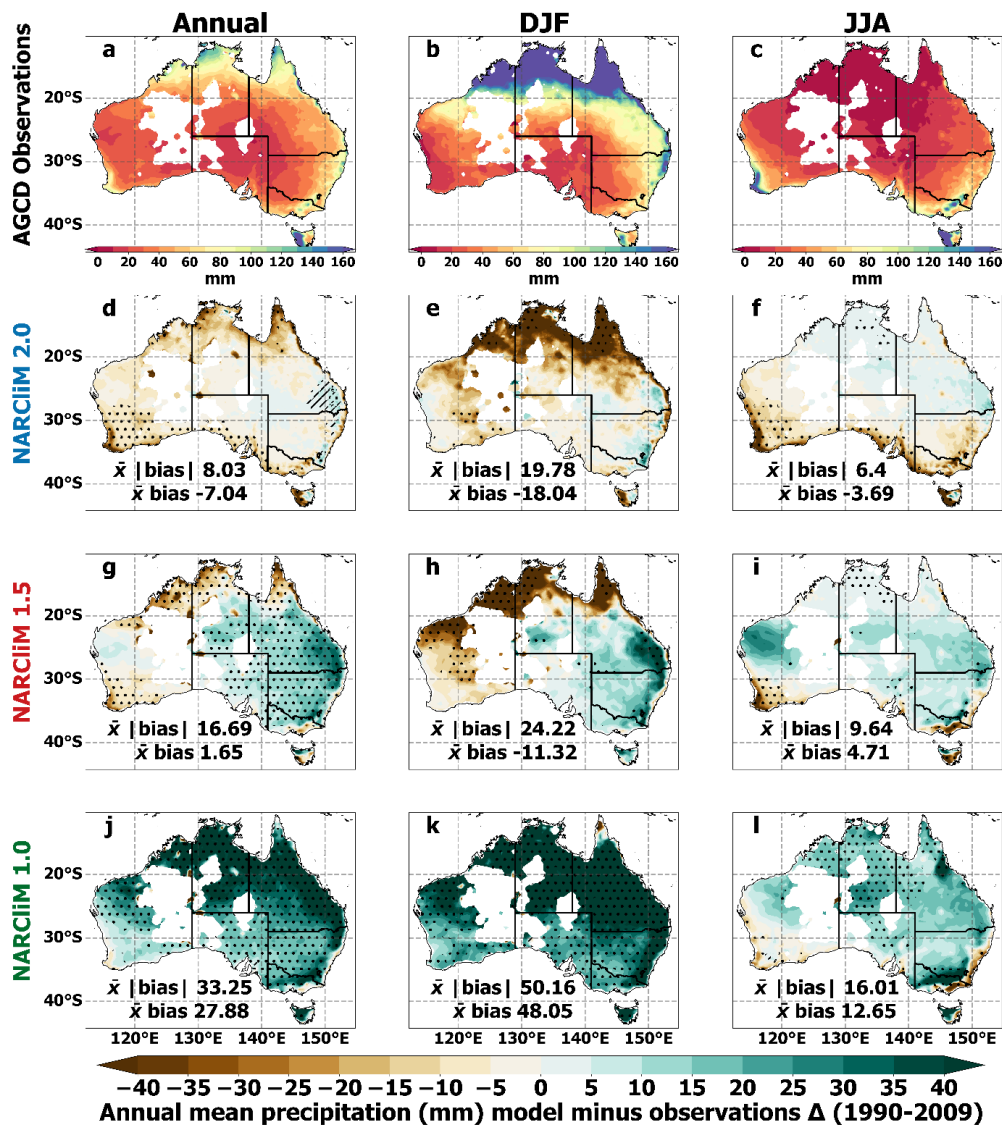
740 mm versus 16.69 mm and 33.25 mm, respectively. Annual mean precipitation biases are particularly  
741 small over eastern regions, often being <5 mm. ~~NARCM2.0~~NARCM 2.0 retains the strong sum-  
742 mertime dry biases for precipitation over northern Australia that are also evident for ~~NAR-~~  
743 ~~CM1.5~~NARCM 1.5 (Figure 11e,h), noting that this region also shows strong warm biases for max-  
744 imum temperature (Figure 9).

745 The individual RCMs comprising ~~NARCM2.0~~NARCM 2.0 show a range of results for an-  
746 nual and seasonal mean precipitation biases (Fig S12-S14). Notably, three of the ten ~~NAR-~~  
747 ~~CM2.0~~NARCM 2.0 RCMs have substantially larger bias magnitudes than their peers at annual and  
748 summer timescales, i.e.i.e., both MPI-ESM1-2-HR-R3 and R5 (absolute biases are 15.53 mm and  
749 22.45 mm for annual mean precipitation, Figure S12g-h) and EC-Earth3-Veg-R5 (Figure S12f; 18.59  
750 mm). Despite EC-Earth3-Veg-R5 being strongly dry-biased, EC-Earth3-Veg-R3 simulates precipita-  
751 tion more accurately i.e.i.e., its mean absolute bias=9.53 mm (Figure S12e). Analogously to ~~NAR-~~  
752 ~~CM2.0~~NARCM 2.0's performances for temperature, R5 is drier than R3. Comparing the ensemble  
753 means of the driving GCMs, the ~~CMPI6~~CMIP6 GCMs are marginally more accurate in simulating  
754 annual mean precipitation than the CMIP5 GCMs (Figure S15). Whilst the CMIP6 ensemble produces  
755 small biases over inland regionsareas, its biases are larger along the east coast.





756

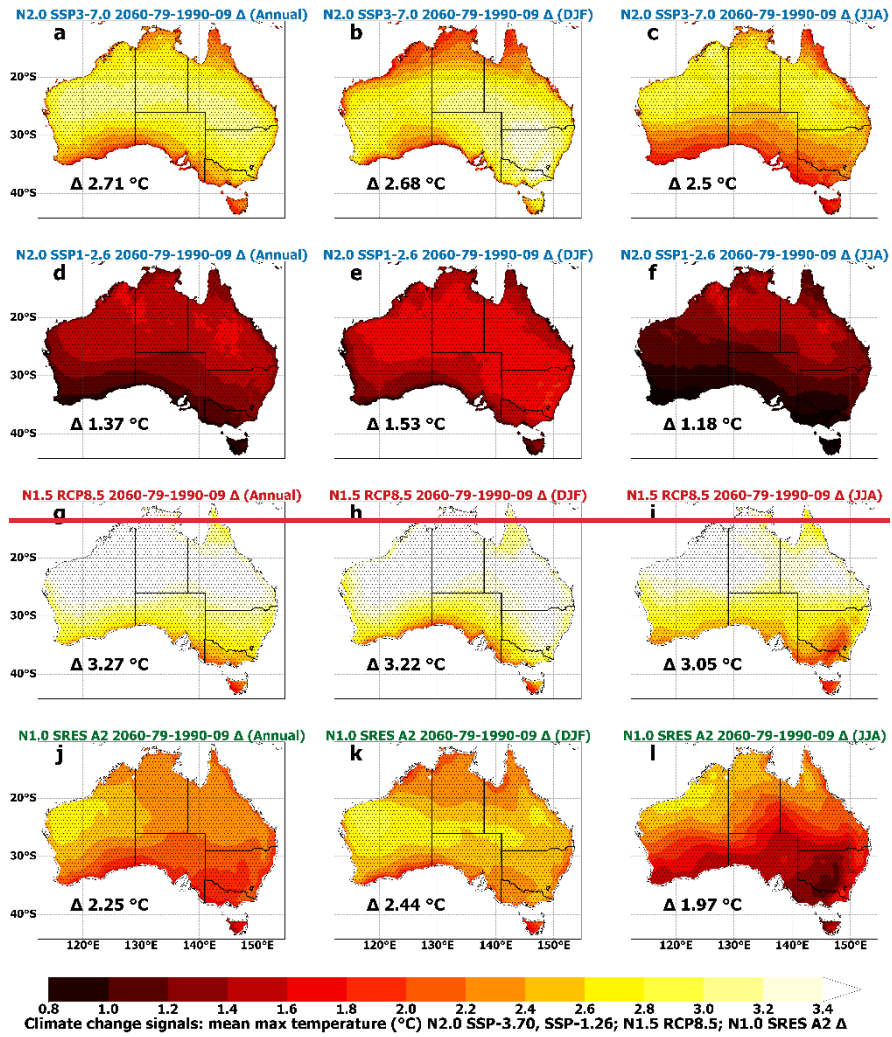


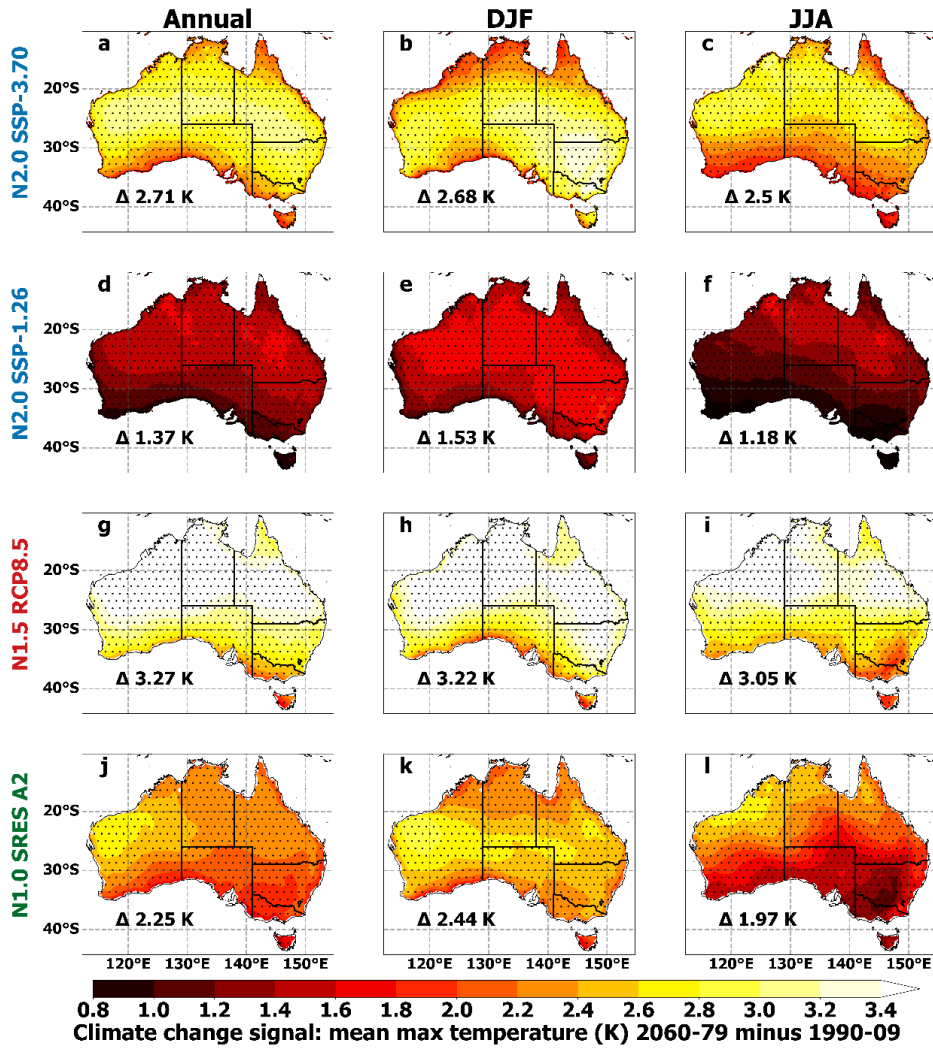
757  
758 Figure 11. As per Figure 9 but for mean precipitation (precip.).

759 **8.7. CORDEX-CMIP6 NARCIIM2.0 NARCIIM 2.0 climate**  
760 **change projections**

761 Dependent on location, the largest maximum temperature projected increases for NAR-  
762 CIIM2.0 NARCIIM 2.0 under SSP3-7.0 are over  $\sim 3\text{ }^{\circ}\text{C}$ , and over  $\sim 1.5\text{ }^{\circ}\text{C}$  under SSP1-2.6 (Figure  
763 12a,d). SSP3-7.0-NARCIIM2.0 NARCIIM 2.0 shows faster warming over inland than coastal regions,  
764 with greater warming across a horizontal band of the continent during annual and summer timescales  
765 (Figure 12a-b). This contrasts with NARCIIM1.5 NARCIIM 1.5 which shows a north-south warming  
766 gradient at annual and seasonal timescales, with its fastest warming rate over northern regions, and  
767 NARCIIM1.0 NARCIIM 1.0 which projects fastest warming over the west (Figure 12). For NAR-  
768 CIIM2.0 NARCIIM 2.0, the tropical north warms faster during the winter dry season than during the  
769 summer wet season under SSP3-7.0, but this is not the case for SSP1-2.6 (Figure 12b-c; e-f). NAR-  
770 CIIM2.0 NARCIIM 2.0 simulations under SSP3-7.0 show less warming than NARCIIM1.5 NARCIIM

771 1.5-RCP8.5, but warmer futures than for ~~NARCIIM1.0~~NARCIIM 1.0-SRES A2, with differences in  
772 the underlying driving GCMs and GHG scenarios likely contributing to these variations in warming.  
773 As per NARCIIM1.x, all ~~NARCIIM2.0~~NARCIIM 2.0 maximum temperature projections are signifi-  
774 cant-agreeing with all RCMs projecting statistically significant temperature increases.

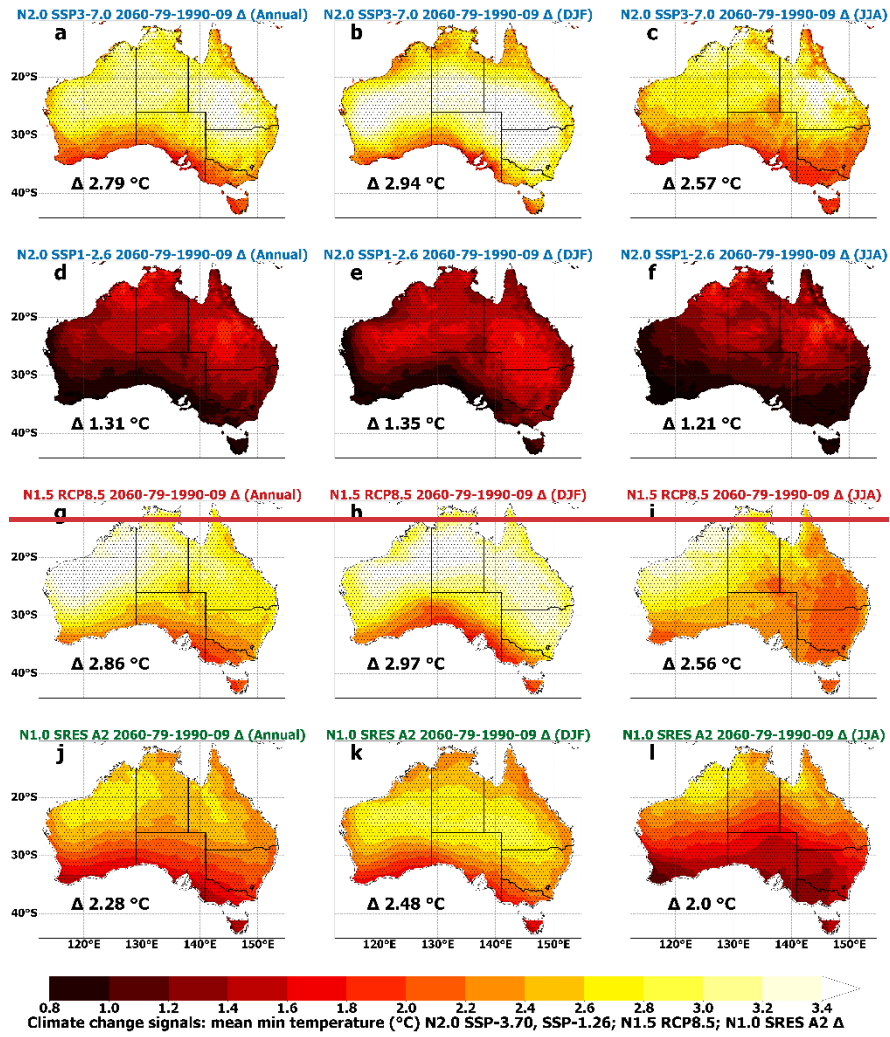


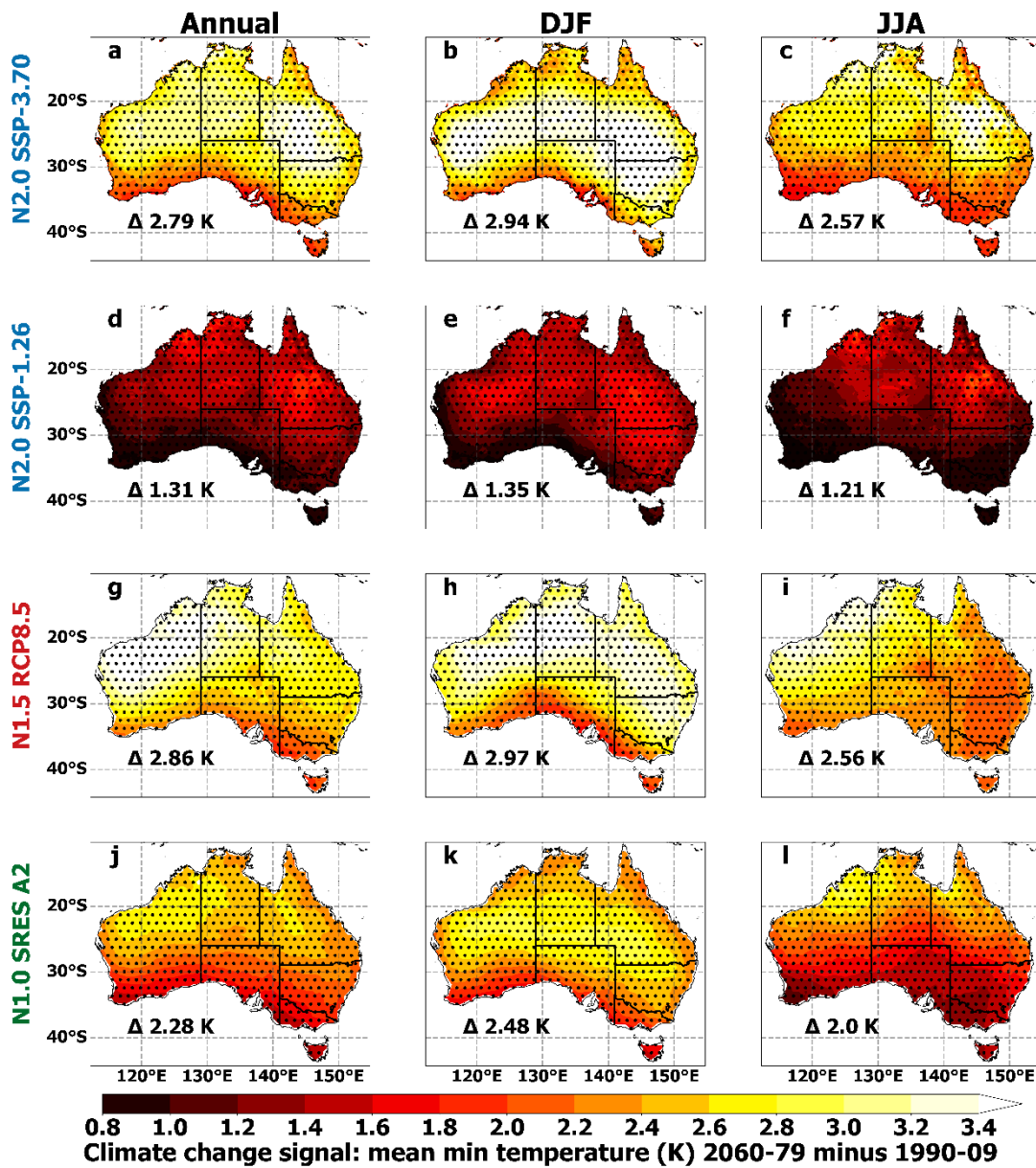


776

777 **Figure 12.** Ensemble mean climate change projections (far future minus present-day) for annual, DJF and JJA  
 778 mean maximum temperatures with significance stippling as per Figure 9.

779 Projected increases in annual mean minimum temperature for [NARClIM2.0](#) exceed  $3.0^{\circ}\text{C}$  over some regions for SSP3-7.0, and  $1.6^{\circ}\text{C}$  for SSP1-2.6 (Figure 13). Under both GHG  
 780 scenarios, at annual and winter timescales warming is fastest over north-east Australia. Conversely,  
 781 [NARClIM1.x](#) minimum temperature future increases are generally largest over northwest or northern  
 782 Australia, though the summertime projection for [NARClIM1.0](#) is an exception (Figure  
 783 13k). As for maximum temperature projections, all RCMs for all [NARClIM](#) generations project statis-  
 784 tically significant increases.  
 785

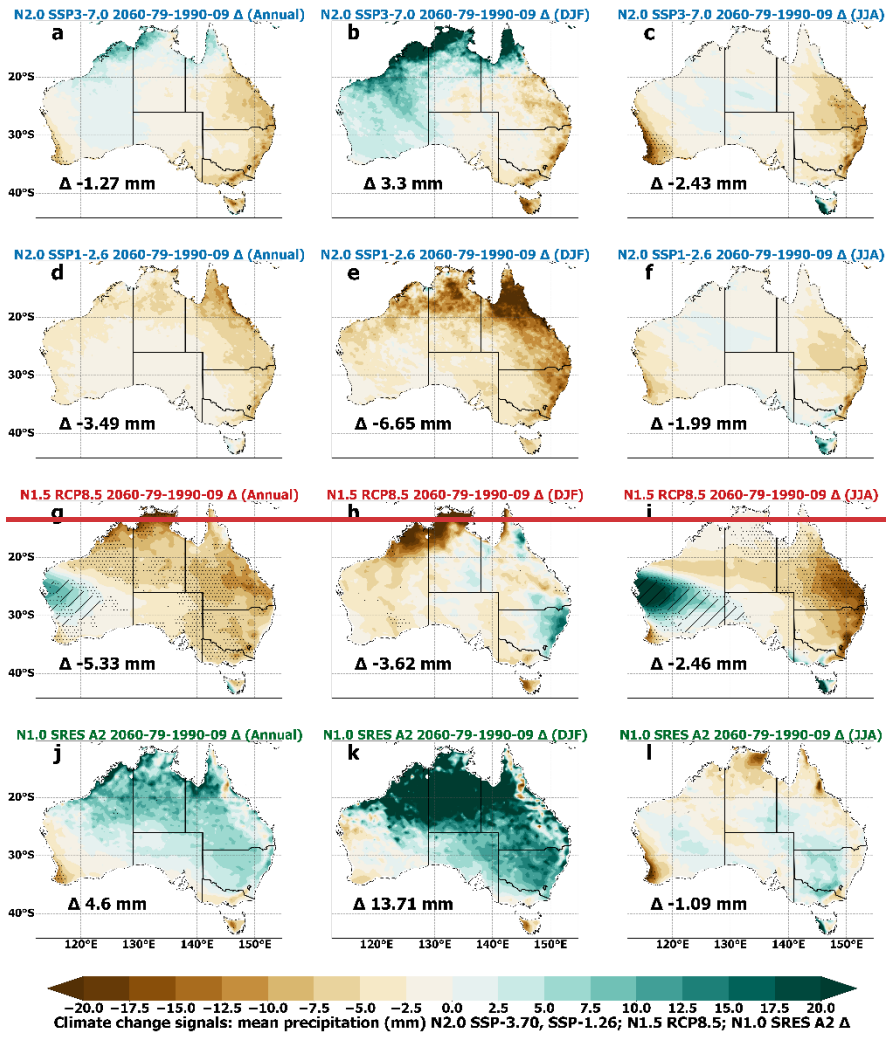




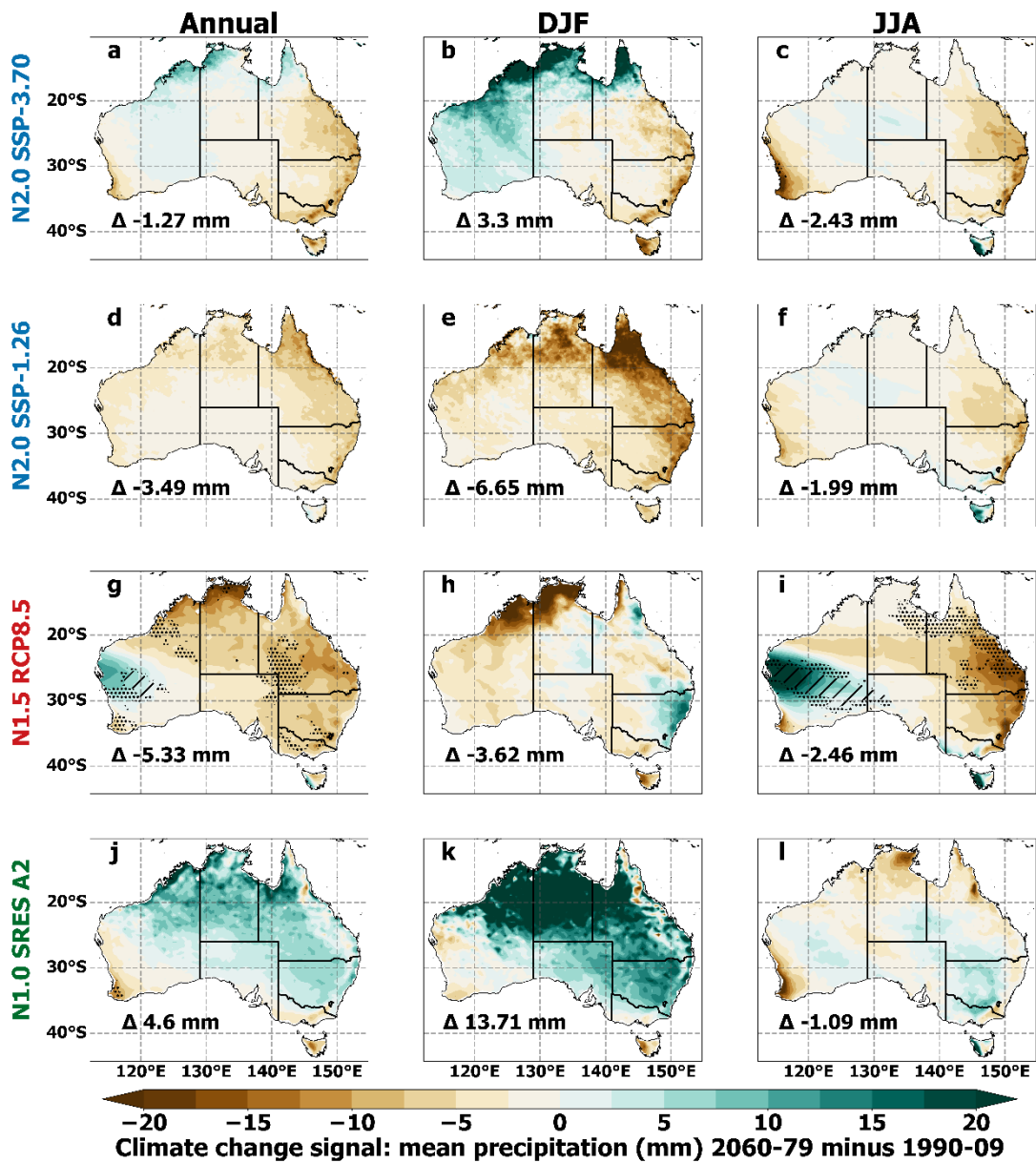
787

788 **Figure 13.** Ensemble mean climate change projections (far future minus present-day) for annual, DJF and JJA  
 789 mean minimum temperatures with significance stippling as per Figure 9.

790 [NARCIIM2.0](#) [NARCIIM 2.0](#) SSP3-7.0 projects a dry future over most of Australia, except for  
 791 wetter futures over northern and western regions, which are largest in magnitude in summer (Figure  
 792 14a-b). In contrast, overall, [NARCIIM2.0](#) [NARCIIM 2.0](#) SSP1-2.6 projects dry changes across most of  
 793 Australia, with the strongest drying over northern Australia during summer (Figure 14e). Similarities  
 794 between [NARCIIM2.0](#) [NARCIIM 2.0](#) projections for the low and high GHG SSPs include faster dry-  
 795 ing over the eastern coastline at all timescales, especially during summer. The wetter futures projected  
 796 by RCMs downscaling SSP3-7.0-GCMs relative to SSP1-2.6 may be partially inherited from the driv-  
 797 ing CMIP6 GCMs, because overall, SSP3-7.0 GCMs show wetter futures than corresponding SSP1-  
 798 2.6 GCMs (Fig. S16).





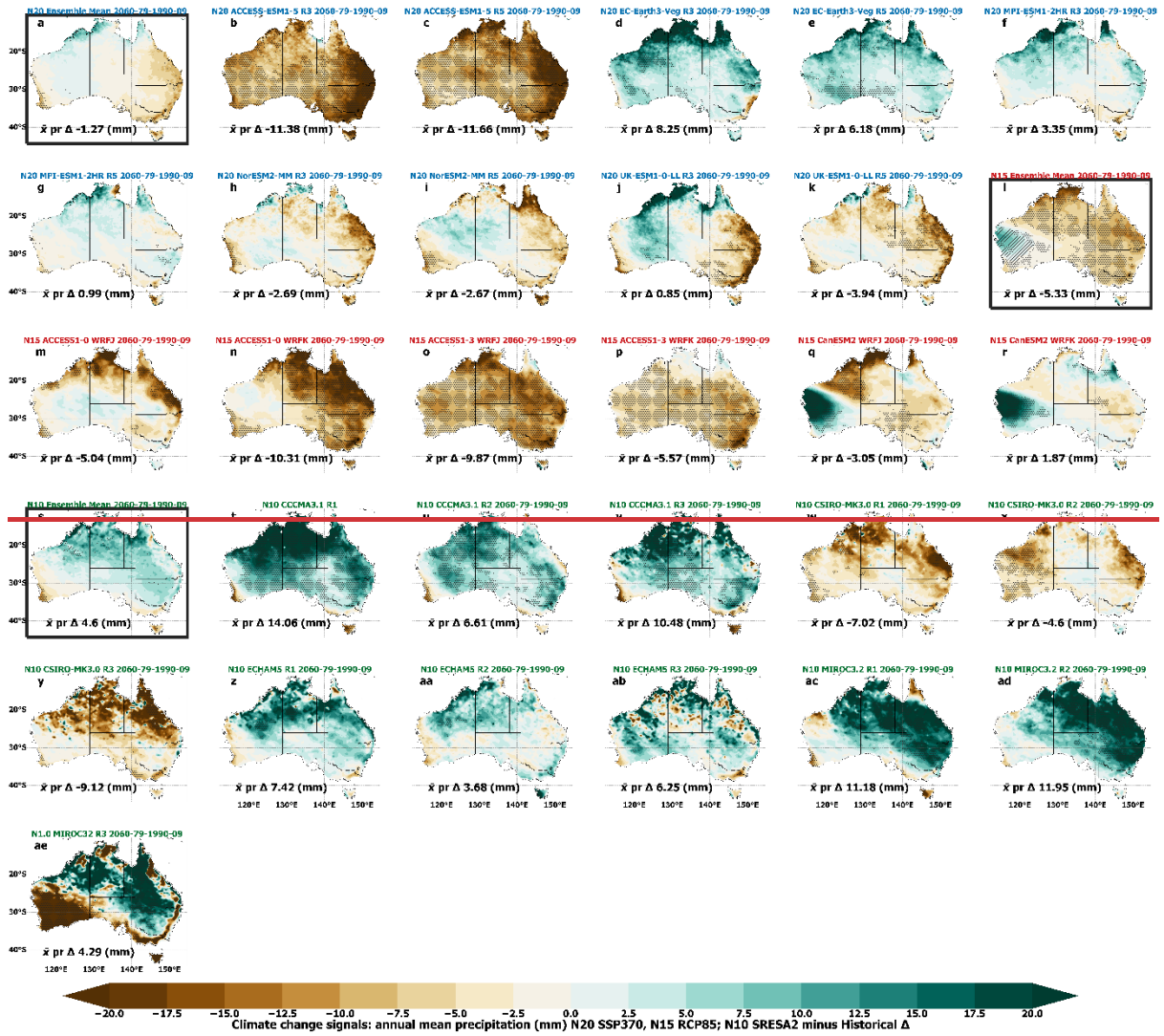


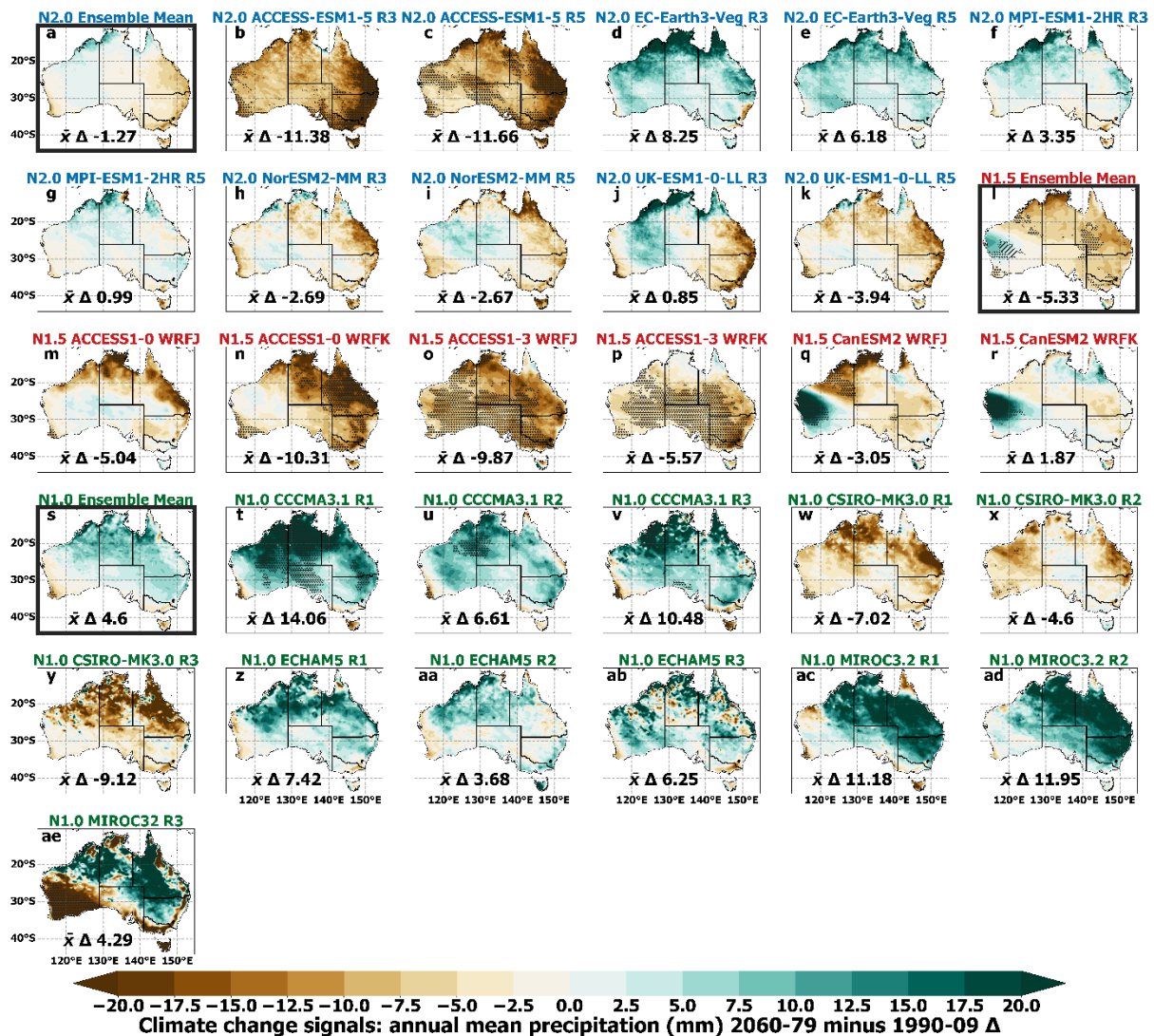
800

801 **Figure 14.** Ensemble mean climate change projections (far future minus present-day) for annual, DJF and JJA  
 802 mean precipitation with significance stippling as per Figure 9.

803 Considering mean precipitation projections for individual ~~NARCIIM2.0~~NARCIIM 2.0 RCMs, in  
 804 some cases, R3 and R5 RCMs produce similar results when downscaling the same GCM. For in-  
 805 stance, ACCESS-ESM-1-5 forced R3 and R5 both show strong projected decreases in annual mean  
 806 precipitation across Australia (Figure 15b-c). In contrast, while UK-ESM1-0-LL R3-R5 both show  
 807 projected decreases in annual mean precipitation over eastern Australia, R3 shows precipitation in-  
 808 creases that are substantially more widespread over western and northern regions relative to R5 (Fig-  
 809 ure 15j-k). Overall, the ~~NARCIIM2.0~~NARCIIM 2.0 ensemble members show a variety of climate  
 810 change signals for precipitation (Figure 15) and temperature (not shown), reflecting the range within  
 811 the larger CMIP6 ensemble (Di Virgilio et al. 2022).

812 There are some key differences between the mean precipitation projections of ~~NAR-~~  
813 ~~CIIM2.0~~NARCIIM 2.0 relative to those of previous NARCIIM generations. For instance, ~~NAR-~~  
814 ~~CIIM1.5~~NARCIIM 1.5 shows stronger reductions in future precipitation over northern and eastern  
815 regions at annual and winter timescales (Figure 14), and these changes are statistically significant over  
816 a few regions, whereas projected changes for ~~NARCIIM2.0~~NARCIIM 2.0 are largely non-significant.  
817 Additionally, ~~NARCIIM2.0~~NARCIIM 2.0 projects marked precipitation decreases along the south-  
818 east coast during summer, while ~~NARCIIM1.5~~NARCIIM 1.5 shows the opposite result (Figure 14).  
819 ~~NARCIIM1.0~~NARCIIM 1.0 generally projects wet futures across larger portions of Australia, espe-  
820 cially at annual and summer timescales.





822

823

824

825

826

827

**Figure 15.** Climate change projections (1990-2009 versus 2060-2079) for annual mean precipitation for NAR-  
 CliM ensemble mean climate change signals (a,l,s) and for individual ensemble members for each generation of  
NARcliM simulation (NARcliM 2.0 under SSP3-7.0, NARcliM 1.5 under RCP8.5 and NARcliM 1.0 under  
SRES A2). Significance stippling as per Figure 9.

## 9.8. Discussion and Summary

828

829

830

831

832

833

834

835

NARcliM regional climate models produce robust climate projections at spatial scales suitable for  
 local-scale climate change analysis and impact decision-making. The third and latest generation of  
 these regional climate models, ~~NARcliM2.0~~ NARcliM 2.0, encompasses several model design ad-  
 vancements over its predecessors. A key aim of this paper is to describe how NARcliM 2.0 differs  
~~from its predecessors and explain the rationale for these design decisions. We also~~ Here, our aims were  
~~to describe the new CORDEX CMIP6 NARcliM2.0 RCM ensemble and its design process, including~~  
~~the model test and evaluation approaches used, and~~ characterise the improvements in model skill in  
 simulating the Australian climate relative to previous NARcliM generations, as well as differences

836 ~~to compare~~ climate projections across NARClIM generations. The next section discusses aspects of  
837 NARClIM2.0 RCM design and parameterisation in relation to previous studies before reviewing dif-  
838 ferences in the model biases and the climate projections of the NARClIM 2.0 versus NARClIM 1.x  
839 RCMs.

## 840 **8.1 ~~NARClIM2.0~~NARClIM 2.0 RCM physics testing**

841 ~~A key aim of this paper is to describe how NARClIM2.0 differs from its predecessors and explain the~~  
842 ~~rationale for these design decisions.~~ In addition to RCM design choices including increased resolu-  
843 tion, and incorporation of convection-permitting modelling and urban physics, a major change for  
844 ~~NARClIM2.0~~NARClIM 2.0 relative to its predecessors is to use new WRF RCM configurations  
845 which are selected via a large suite of physics tests. RCM performance evaluations for the ~~NAR-~~  
846 ~~ClIM2.0~~NARClIM 2.0 RCM physics testing focused on the 4 km resolution convection-permitting  
847 domain which does not use a cumulus physics parameterisation. Notably, the 7 candidate shortlisted  
848 RCMs from the N=78 physics test ensemble used three different cumulus parameterisations for their  
849 outer domains, with 4 RCMs using BMJ, 2 RCMs using Tiedtke, and 1 using Kain-Fritsch. This indi-  
850 cates that differences in the outer domain boundary conditions have key influences on the RCM per-  
851 formances in the convection-permitting domain. ~~Notably, the 7 ‘candidate’ shortlisted RCMs from the~~  
852 ~~N=78 physics test ensemble used three different cumulus parameterisations for their outer domains,~~  
853 ~~with 4 RCMs using BMJ, 2 RCMs using Tiedtke, and 1 using Kain-Fritsch. This indicates that differ-~~  
854 ~~ences in the outer domain boundary conditions have key influences on the RCM performances in the~~  
855 ~~convection-permitting domain.~~

856 ~~The u~~Using of the Noah-MP LSM in the ~~NARClIM2.0~~NARClIM 2.0 RCM physics tests  
857 conferred overall RCM skill improvements relative to ~~the test Phase I~~ RCMs using the Noah-Unified  
858 LSM, especially in terms of the simulation of temperature. Although using Noah-MP also improved  
859 the simulation of precipitation in some respects, these improvements were smaller relative to the gains  
860 for temperature, and improvements were mainly located over coastal regions. The developers of No-  
861 ah-MP suggest that some limitations in the Noah-Unified LSM have been modified to better represent  
862 several parameters. These include surface layer radiation balances, snow depth, soil moisture and heat  
863 fluxes, leaf area-rainfall interaction, vegetation and canopy temperature distinction, drainage of soil,  
864 and runoff.

865 The developers of Noah-MP suggest that some limitations in the Noah-Unified LSM have  
866 been modified to better represent several parameters. These include surface layer radiation balances,  
867 snow depth, soil moisture and heat fluxes, leaf area-rainfall interaction, vegetation and canopy tem-  
868 perature distinction, drainage of soil, and runoff.

869 In the NARClIM2.0 physics testing, improvements in RCM skill were evident for Noah-MP  
870 with default settings. Activating specific parameterisations for this LSM (i.e. dynamic vegetation cov-

871 er and surface runoff-simple groundwater) delivered comparatively smaller gains in RCM perfor-  
872 mances. Some previous studies have found no overall benefit of using Noah-MP with default settings.  
873 For instance, In the NARClIM2.0 physics testing, improvements in RCM skill were evident for Noah-  
874 MP with default settings. Activating specific parameterisations for this LSM (i.e. dynamic vegetation  
875 cover and surface runoff-simple groundwater) delivered comparatively smaller gains in RCM perfor-  
876 mances. Some previous studies have found no overall benefit of using Noah-MP with default settings.  
877 For instance, Imran et al. (2018) conducted an evaluation of WRF coupled with a variety of LSMs  
878 including Noah-MP using its default settings. ~~Their focus was on~~ simulated short-duration (~3-  
879 day) heatwaves in Melbourne, Australia. ~~They and~~ observed larger temperature biases using Noah-  
880 MP relative to RCMs using Noah-Unified and CLM4.0. However, their focus on specific short dura-  
881 tion heatwave events ~~of short duration~~ over one urban area was not intended as a comprehensive eval-  
882 uation of Noah-MP's performance ~~over longer timescales. It is also important to consider that~~ Addi-  
883 tionally, several physics schemes used by these authors differed to those used in the ~~NAR-~~  
884 ~~ClIM2.0~~ NARClIM 2.0 physics testing, ~~i.e.i.e.,~~ they used: PBL=MYJ; microphysics=Thompson; cum-  
885 ulus=Grell3D; radiation=RRTMG/RRTMG. ~~The o~~ Only similarities between these settings and  
886 those of the NARClIM2.0 physics testing are the use of Thompson microphysics and RRTMG radia-  
887 tion are used in the NARClIM 2.0 physics testing. WRF and Noah-MP versions also differed, ~~i.e.i.e.,~~  
888 Imran et al. used WRF3.6.1 and a Noah-MP version prior to 3.7, whereas ~~NARClIM2.0~~ NARClIM 2.0  
889 uses WRF4.1.2 and Noah-MP version 4.1. Additionally, there are also several studies that have re-  
890 ported benefits of using Noah-MP with default parameters relative to other LSMs for other regions  
891 globally e.g. Chen et al. (2014b), Chen et al. (2014a) and Salamanca et al. (2018).

892 In an assessment of the performances of several WRF LSMs for Sub-Saharan Africa,  
893 Glotfelty et al. (2021) noted deficiencies in the simulation of land use and land cover change  
894 (LULCC) parameters such as surface albedo by Noah-MP. Despite these deficiencies, the spatial pat-  
895 terns and magnitudes of temperature and precipitation were well represented by Noah-MP. However,  
896 the land surface parameter errors impacted the magnitude and sign of LULCC-induced changes in  
897 temperature and precipitation. These deficiencies were linked to substantial underestimations of sur-  
898 face albedo in arid areas due to inaccurate soil albedo treatments by Noah-MP. Moreover, errors in  
899 Noah-MP's LAI profiles may occur because it was developed principally for application in Northern  
900 Hemisphere mid-latitudes. It is possible that modifying/tuning Noah-MP to specific aspects of the  
901 Australian context would yield performance benefits for follow-up dynamical downscaling. Overall,  
902 these authors concluded that "Noah-MP is least flawed of the [WRF] default LSMs". Additionally,  
903 there are also several studies that have reported benefits of using Noah-MP with default parameters  
904 relative to other LSMs e.g. Chen et al. (2014b), Chen et al. (2014a) and Salamanca et al. (2018).

905 The ~~NARClIM2.0~~ NARClIM 2.0 physics testing found that the optimal LSM configuration for  
906 simulation of minimum temperature used Noah-MP with dynamic vegetation cover activated, even  
907 though the performance gain relative to Noah-MP with default settings was small. Constantinidou et

908 al. (2020) ran WRF coupled with four LSMs (Noah-Unified, Noah-MP, CLM and, Rapid Update Cy-  
909 cle) over the Middle East North Africa CORDEX domain. The ~~vir study~~ compared the performance of  
910 Noah-MP with dynamic vegetation cover turned on and off. ~~They and showed found~~ that air and land  
911 temperatures were best simulated using Noah-MP with dynamic vegetation cover activated.

912 ~~Overall, Noah MP performed well in the NARClIM2.0 physics tests, conferring some clear~~  
913 ~~advantages over RCMs using Noah Unified. However, given the nature of its development and per-~~  
914 ~~formance characteristics, it may be more suited to application over the temperate regions of Australia~~  
915 ~~rather than the semi-arid interior.~~

916 In terms of ~~PBL other NARClIM2.0 RCM~~ parameterisations, ~~focusing on PBL~~, by the com-  
917 pletion of Phase I physics testing, only 3 of 12 RCMs shortlisted for further testing use the YSU  
918 scheme. By the completion of Phase II testing, all remaining RCMs using YSU are discarded, with  
919 only RCMs using PBL schemes other than YSU remaining (~~i.e.i.e.~~, ACM2 and MYNN2). YSU PBL  
920 is a first-order closure scheme that expresses turbulent mixing via mean variables rather than prognos-  
921 tic variables (Hong et al., 2006). It is classed as a 'non-local' scheme because it estimates turbulent  
922 mixing by small-scale eddies as well as representing transport caused by convective large eddies. Two  
923 previous studies evaluating convection permitting WRF simulations using different parameterisations  
924 that included YSU for the PBL scheme found that, relative to other PBL schemes, YSU produced the  
925 highest bias for simulated precipitation (Huang et al., 2023; Nuryanto et al., 2019). However, these  
926 studies focused on different regions globally and used various experimental setups that are not direct-  
927 ly comparable to those used here. Hence, a separate study investigating sensitivities of the ~~NAR-~~  
928 ~~ClIM2.0NARClIM 2.0~~ RCMs to the different PBL schemes is currently underway.

## 929 **8.2 CORDEX-CMIP6 ~~NARClIM2.0NARClIM 2.0~~: historical evaluation** 930 **and climate change projections**

931 We characterised the improvements conferred by ~~NARClIM2.0NARClIM 2.0~~ over its predecessors in  
932 simulating the present-day Australian climate. ~~NARClIM2.0NARClIM 2.0~~ simulates mean maximum  
933 temperature and precipitation more accurately than NARClIM1.x. Specifically, NARClIM1.x has  
934 strong maximum temperature cold biases which are in keeping with other downscaling projects of the  
935 CMIP3-CMIP5 eras, e.g., (Andrys et al., 2016; Evans et al., 2020b), but these are substantially re-  
936 duced in ~~NARClIM2.0NARClIM 2.0~~. A contributing cause of CMIP5-forced RCM cold biases of  
937 maximum temperature is their overestimation of precipitation (Evans et al., 2020). This relationship  
938 was also noted in ERA-Interim forced RCMs of this ~~same~~ modelling era (Di Virgilio et al. 2019). In  
939 ~~NARClIM2.0NARClIM 2.0~~, the widespread wet biases that characterise the NARClIM1.x RCMs are  
940 ~~greatly~~ reduced in magnitude. ~~NARClIM2.0NARClIM 2.0~~ produces smaller wet biases over eastern  
941 Australia, and smaller dry biases elsewhere, except for Australia's tropical north. This marked reduc-  
942 tion in wet bias magnitudes is ~~aone~~ plausible contributing factor for the reduction in maximum tem-

943 perature cold bias for the ~~NARClIM2.0~~NARClIM 2.0 RCMs. The CMIP6 and CMIP5 GCMs used to  
944 drive ~~NARClIM2.0~~NARClIM 2.0 and 1.5 RCMs generally show similar magnitudes of maximum  
945 temperature cold bias. This suggests that the underlying nature of the CMIP6 driving data is not a  
946 principal factor underlying the observed improvements for ~~NARClIM2.0~~NARClIM 2.0's simulation  
947 of maximum temperature. In fact, the RCMs appear to have a substantial influence on the reduced  
948 maximum temperature biases.

949 That ~~NARClIM2.0~~NARClIM 2.0 underestimates precipitation over tropical northern Australia  
950 during the wet season (summer) to a similar degree of magnitude to the ~~NARClIM1.5~~NARClIM 1.5  
951 RCMs indicates that the newer models still struggle to accurately capture the strength of the Australi-  
952 an monsoon. That NARClIM1.x strongly overestimates precipitation over south-eastern Australia,  
953 whereas wet biases over this region are reduced for ~~NARClIM2.0~~NARClIM 2.0 indicates that the  
954 newer models may confer an improved simulation of broad-scale processes associated with synoptic-  
955 scale systems interacting with the extratropical storm track over Australia (Grose et al., 2019).

956 ~~In terms of whether~~The extent to which NARClIM2.0's improved simulation of precipitation  
957 ~~is-might be~~ attributable to ~~theits~~ driving data warrants consideration. Overall, the CMIP6 GCMs  
958 used to drive NARClIM 2.0 show marginally reduced wet biases versus the CMIP5 GCMs used for  
959 NARClIM1.5 (e.g. area-averaged ensemble mean absolute biases are 7.13 mm and 8.89 mm, respec-  
960 tively; Supporting Information Figure S15). This suggests that the underlying nature of the CMIP6  
961 driving data ~~is-might not be~~ the principal factor underlying the observed improvements for NARClIM  
962 2.0's simulation of mean precipitation. Conversely, in terms of RCM design features, the use of the  
963 Noah-MP LSM in the NARClIM 2.0 RCM physics tests conferred overall RCM skill improvements  
964 relative to RCMs using the Noah-Unified LSM for both mean precipitation and mean maximum tem-  
965 perature. As noted above, ~~T~~the developers of Noah-MP suggest that some features of the Noah-  
966 Unified LSM have been modified to better represent several parameters ~~such as soil moisture and heat~~  
967 ~~fluxes, leaf area rainfall interaction, vegetation and canopy temperature distinction, drainage of soil,~~  
968 ~~and runoff~~. The production NARClIM2.0 RCMs ~~forced with CMIP6 GCMs~~ used Noah-MP, whereas  
969 NARClIM1.x RCMs used Noah-Unified. Given these performance improvements observed for RCMs  
970 using Noah-MP versus ~~RCMs~~ using Noah-Unified, it is plausible that the different newer land surface  
971 scheme LSMs (i.e. Noah MP for NARClIM 2.0 versus Noah Unified for NARClIM 1.x) play a role  
972 ~~in~~contributes to the improved NARClIM2.0 RCM-skill in simulating ~~mean~~ precipitation and maxi-  
973 mum temperature, for instance, via changing the land surface feedback (via soil moisture) to the simu-  
974 lation of precipitation. ~~However,~~ ~~t~~This possibility requires more extensive investigation via future  
975 studies.

976 More generally, ~~in this scope of the present study,~~ ~~the scope~~ was to focus on an initial "first-  
977 order" evaluation of mean precipitation rather than extremes of precipitation. However, clearly valua-  
978 ble research can now be undertaken into evaluating the skill of NARClIM 2.0 in simulating extreme  
979 precipitation, subdaily precipitation, etc, using NARClIM 2.0 20 km and 4 km data, noting these data



980 are now publicly available. A good avenue for further research is to assess the potential added value  
981 in simulating extreme and subdaily precipitation at convection permitting scale versus the convection-  
982 parameterised 20 km data. Several previous studies have confirmed that convection-permitting resolu-  
983 tion models can improve the simulation of daily and sub-daily rainfall extremes (Xie et al., 2024;  
984 Cannon and Innocenti, 2019; Kendon et al., 2017).

985 ~~\_\_\_\_\_~~ ~~NARCM2.0~~ NARCM 2.0 RCMs overestimate minimum temperatures across Aus-  
986 tralia, and these biases are larger relative to ~~NARCM1.5~~ NARCM 1.5 but comparable to those of  
987 ~~NARCM1.0~~ NARCM 1.0. The CMIP6 GCMs used to force ~~NARCM2.0~~ NARCM 2.0 show sub-  
988 stantially stronger warm biases for minimum temperature than the CMIP5 GCMs used for ~~NAR-~~  
989 ~~CM1.5~~ NARCM 1.5. This suggests that the increased warm bias for minimum temperature in  
990 ~~NARCM2.0~~ NARCM 2.0-RCMs ~~is~~ could be partially inherited from the driving GCMs. However,  
991 ~~as noted above, the~~ Noah-MP's LSM simulation of factors such as LAI and other aspects of vegeta-  
992 tion as well as surface albedo in semi-arid and arid areas has been shown to have deficiencies  
993 (Glotfelty et al., 2021). These issues may contribute to some of the biases shown by the ~~NAR-~~  
994 ~~CM2.0~~ NARCM 2.0 RCMs. Moreover, the ~~NARCM2.0~~ NARCM 2.0 ensemble mean reduces the  
995 overall minimum temperature bias of the CMIP6 GCM ensemble by almost half, attesting to the add-  
996 ed value conferred by the ~~NARCM2.0~~ NARCM 2.0 RCMs with respect to near-surface temperature  
997 variables.

998 Consideration of observational uncertainty is warranted. We have evaluated NARCM RCM  
999 skill via comparison with AGCD observations. Whilst AGCD are a high quality gridded observational  
1000 data set, like any set of observations, they contain errors and uncertainties. Consequently, the out-  
1001 comes of our evaluations depend on both the models being evaluated and the AGCD observational  
1002 dataset. This is clearly a broader issue that applies to any model evaluation versus observations. Un-  
1003 certainties in AGCD for temperature and precipitation arise from sparse station coverage in some lo-  
1004 cations, especially in remote areas, and interpolation errors in generating gridded data. More specifi-  
1005 cally, temperature uncertainties include urban heat island effects, inhomogeneities in observation rec-  
1006 ords, and elevation differences. Precipitation uncertainties involve underestimation of extremes, rain  
1007 gauge measurement errors, and challenges in representing complex terrain. For our purposes, the  
1008 question of how much of a model bias of ~0.5 K is due to the model errors versus the observational  
1009 uncertainty cannot be currently quantified, because the models are evaluated against this single obser-  
1010 vational dataset. This leaves the observational uncertainty as implicitly included in our results. In the  
1011 future observational uncertainty could be explicitly considered using a method like the Observation  
1012 Range Adjusted (ORA) statistics (Evans and Imran, 2024).

### 1013 8.3 CORDEX-CMIP6 NARClIM 2.0 climate change projections

1014 In terms of ~~NARClIM2.0~~NARClIM 2.0 future climate projections, major changes between NARClIM  
1015 generations such as differences in GHG scenarios mean that ~~NARClIM2.0~~NARClIM 2.0 projected  
1016 temperature changes differ in some respects to those of its predecessors. Overall, as is expected, pro-  
1017 jected warming is less intense in ~~NARClIM2.0~~NARClIM 2.0 under SSP3-7.0 than for ~~NAR-~~  
1018 ~~ClIM1.5~~NARClIM 1.5 under RCP8.5. Other differences in the projections between NARClIM genera-  
1019 tions require further investigation in order to explain, such as ~~NARClIM1.5~~NARClIM 1.5's latitudinal  
1020 warming gradient for maximum temperature that contrasts with ~~NARClIM2.0~~NARClIM 2.0's band of  
1021 faster warming over central Australia relative to northern and southern regions. Irrespective of these  
1022 differences, all three NARClIM ensembles show widespread statistically significant-agreeing results  
1023 for warming projections.

1024 Precipitation projections for the different NARClIM generations show some key similarities,  
1025 such as reductions in mean annual precipitation over eastern Australia for ~~NARClIM2.0~~NARClIM 2.0  
1026 and ~~NARClIM1.5~~NARClIM 1.5, though a difference is that these are statistically significant over  
1027 some areas only for ~~NARClIM1.5~~NARClIM 1.5. The ~~NARClIM2.0~~NARClIM 2.0-SSP3-7.0 and  
1028 SSP1-2.6 ensembles differ in that the former generally projects wet changes over northern and west-  
1029 ern Australia, whereas the latter is generally dry, results that appear partially traceable to the underly-  
1030 ing driving CMIP6-SSP data (Supporting Information Figure S16). ~~Other notable differences are that~~  
1031 ~~some NARClIM2.0 RCMs produce very similar precipitation projections for certain GCM-RCM~~  
1032 ~~combinations, such as for~~

1033 Some ~~NARClIM2.0~~NARClIM 2.0 RCMs produce very similar precipitation projections for  
1034 certain GCM-RCM combinations. Notably, ~~-ACCESS-ESM-1-5 forced -R3 versus and~~ R5 under  
1035 SSP3-7.0 both produce ~~(i.e.~~ widespread dry projections that are substantially drier than other NAR-  
1036 ClIM 2.0 models for both RCMs). This GCM projects very dry futures across Australia (Di Virgilio et  
1037 al., 2022), so this result in the R3 and R5 RCMs could be largely inherited from the driving data.  
1038 There are 40 realisations for ACCESS-ESM1-5, but only realisation 6 provides sub-daily outputs that  
1039 can be used in dynamical downscaling using WRF. This realisation simulates a particularly dry pro-  
1040 jection over Australia, especially for eastern Australia, making it a useful "stress test" case. In terms  
1041 of GCM skill versus observations, globally, this GCM is dry biased over a few regions owing to a lo-  
1042 cation bias with the Inter-tropical Convergence Zone (Rashid et al., 2022; Ziehn et al., 2020).

1043 Conversely, iIn other instances, there are marked divergences between the NARClIM 2.0 R3  
1044 versus R5 precipitation projections when forced with the same GCM. An example is for instance,  
1045 UK-ESM-1-0-LL under SSP3-7.0 where R3 projects stronger precipitation increases that are more  
1046 geographically widespread relative to R5. This raises the question of varying sources of uncertainty in  
1047 the climate projections, ~~i.e.~~ to what extent these are attributable to GCMs versus RCMs, as well as  
1048 other factors.

## 1049 8.4 Summary

1050 In summary, the CORDEX-CMIP6 ~~NARClIM2.0~~NARClIM 2.0 regional climate projections are a 10-  
1051 member ensemble comprising two configurations of the WRF RCM dynamically downscaling five  
1052 GCMs under three SSPs at 20 km resolution over CORDEX-Australasia and at 4 km convection-  
1053 permitting resolution over south-east Australia. In addition to several high-level model design chang-  
1054 es, e.g., increased spatial resolution, a large (N=78) RCM-physics test suite is evaluated to select two  
1055 new WRF RCM configurations for CMIP6-forced ~~NARClIM2.0~~NARClIM 2.0 climate projections.

~~1056 Due to resource constraints and the aim to test a large number of RCM physics parameterisations, the  
1057 NARClIM2.0 physics tests are performed for a single year. This is one reason why the final selection  
1058 of two production grade RCMs for the CMIP6 NARClIM2.0 runs is based on the CORDEX ERA5-  
1059 forced 42-year evaluation simulations.~~ The ~~NARClIM2.0~~NARClIM 2.0 physics tests identified RCM  
1060 configurations that generally performed well in simulating the recent Australian climate over south-  
1061 east Australia. A key finding was that WRF RCMs using the Noah-MP LSM generally out-performed  
1062 RCMs using other ~~WRF~~-LSMs in representing regional climate. Despite the overall performance  
1063 gains evident for RCMs using Noah-MP, ~~RCM these improvements skill was~~ ere superior over ~~the~~  
1064 temperate/coastal regions of southeast Australia, relative to the semi-arid interior. These performance  
1065 characteristics might be linked to Noah-MP's development being focused on Northern Hemisphere  
1066 mid-latitudes, including assumptions such as accounting for differences in seasonality in the Northern  
1067 versus Southern Hemispheres by shifting the Northern Hemisphere LAI profiles by 6 months. For the  
1068 southeast Australian context, noting its distinctive coastal dry-sclerophyll and expansive inland grass-  
1069 land biomes, such assumptions might lead to discontinuities in quantities such as LAI. ~~Hence, future  
1070 investigation into processes such as land surface coupling in NARClIM2.0 RCMs is warranted. Given  
1071 the geographic focus of Noah-MP's development, as well as its performance characteristics, it may be  
1072 more suited to application over the temperate regions of Australia rather than the semi-arid interior. It  
1073 is also possible that modifying/tuning Noah-MP to specific aspects of the Australian context would  
1074 yield performance benefits for follow-up dynamical downscaling.~~

1075 Overall, the CMIP6-~~NARClIM2.0~~NARClIM 2.0 ensemble produces a good representation of  
1076 recent mean climate that in several key respects improves upon the model skill of earlier NARClIM  
1077 generations. This study provides a foundation for more detailed investigations of the model biases and  
1078 future climate changes described here, including process-focused studies exploring their mechanisms.  
1079 CORDEX-CMIP6 ~~NARClIM2.0~~NARClIM 2.0 RCM data provide valuable resources to investigate  
1080 projected climate changes, their impacts on societies and natural systems, and potential climate  
1081 change mitigation and adaptation actions for the CORDEX-Australasia region.

## 1082 **9. Code Availability**

1083 [A frozen version of the source code for the Weather Research and Forecasting \(WRF\) version 4.1.2](#)  
1084 [used in this study, as well as the configuration files for the simulations, is available on Zenodo at:](#)  
1085 <https://doi.org/10.5281/zenodo.11184830>~~The Weather Research and Forecasting (WRF) version 4.1.2~~  
1086 ~~used in this study is freely available from:~~ <https://github.com/coecms/WRF/tree/V4.1.2>. ~~A static copy~~  
1087 ~~of all scripts used for this study can be found at:~~ [https://bitbucket.org/oeheas/narelim2-](https://bitbucket.org/oeheas/narelim2-0-design-and-evaluation-2024-support-materials/src/main/)  
1088 ~~0-design-and-evaluation-2024-support-materials/src/main/~~

## 1089 **10. Data Availability**

1090 Data for the ~~NARCIIM2.0~~[NARCIIM 2.0](#) CMIP6-forced R3 and R5 RCMs are being made available  
1091 via [National Computing Infrastructure](#) (NCI). WRF namelist settings for the ~~NARCIIM2.0~~[NARCIIM](#)  
1092 [2.0](#) CMIP6-forced R3 and R5 RCMs are shown in Supplementary Material Figure S1 and are also  
1093 available at: <https://doi.org/10.5281/zenodo.11184830>. Data ~~NARCIIM1.5~~[NARCIIM 1.5](#) RCMs are  
1094 available via the [New South Wales Climate Data Portal](#) and [CORDEX-DKRZ](#). Data for  
1095 ~~NARCIIM1.0~~[NARCIIM 1.0](#) RCMs are available via the [New South Wales Climate Data Portal](#).  
1096 CMIP6 GCM data are available via the [Earth System Grid Federation](#).

## 1097 **11. Author Contribution**

1098 GDV and JPE designed the models and the simulations. FJ, ET, and CT setup the models and  
1099 conducted the model simulations with contributions from JPE, JK, JA, DC, CR, SW, YL, MER, RG  
1100 and JL. GDV prepared the manuscript with contributions from all co-authors.

## 1101 **12. Competing Interests**

1102 The authors declare that they have no conflict of interest, noting that JK is a Topic Editor of  
1103 Geoscientific Model Development.

## 1104 **13. Funding**

1105 This research was supported by the New South Wales Department of Climate Change, Energy, the  
1106 Environment and Water as part of the ~~NARCIIM2.0~~[NARCIIM 2.0](#) dynamical downscaling project  
1107 contributing to CORDEX Australasia. Funding was provided by the NSW Climate Change Fund for  
1108 NSW and Australia Regional Climate Modelling (NARCIIM) Project. This research was undertaken  
1109 with the assistance of resources and services from the National Computational Infrastructure (NCI),  
1110 which is supported by the Australian Government.

## 1111 **14. References**

- 1112 Andrys, J., Lyons, T. J., and Kala, J.: Evaluation of a WRF ensemble using GCM boundary conditions  
1113 to quantify mean and extreme climate for the southwest of Western Australia (1970–1999),  
1114 International Journal of Climatology, 36, 4406-4424, <https://doi.org/10.1002/joc.4641>, 2016.
- 1115 Australian Bureau of Statistics.: Regional population, Online at:  
1116 <https://www.abs.gov.au/statistics/people/population/regional-population/latest-release>, 2024.
- 1117 Bjordal, J., Storelvmo, T., Alterskjaer, K., and Carlsen, T.: Equilibrium climate sensitivity above 5  
1118 degrees C plausible due to state-dependent cloud feedback, Nat. Geosci., 13, 718-+,  
1119 10.1038/s41561-020-00649-1, 2020.
- 1120 Bureau of Meteorology.: Annual climate statement 2016, 2017.
- 1121 Cannon, A. J. and Innocenti, S.: Projected intensification of sub-daily and daily rainfall extremes in  
1122 convection-permitting climate model simulations over North America: implications for future  
1123 intensity–duration–frequency curves, Nat. Hazards Earth Syst. Sci., 19, 421-440,  
1124 10.5194/nhess-19-421-2019, 2019.
- 1125 Chen, F., Liu, C. H., Dudhia, J., and Chen, M.: A sensitivity study of high-resolution regional climate  
1126 simulations to three land surface models over the western United States, Journal of  
1127 Geophysical Research-Atmospheres, 119, 7271-7291, 10.1002/2014jd021827, 2014a.
- 1128 Chen, F., Barlage, M., Tewari, M., Rasmussen, R., Jin, J. M., Lettenmaier, D., Livneh, B., Lin, C. Y.,  
1129 Miguez-Macho, G., Niu, G. Y., Wen, L. J., and Yang, Z. L.: Modeling seasonal snowpack  
1130 evolution in the complex terrain and forested Colorado Headwaters region: A model  
1131 intercomparison study, Journal of Geophysical Research-Atmospheres, 119, 13795-13819,  
1132 10.1002/2014jd022167, 2014b.
- 1133 Chou, M. D., Suarez, M. J., Liang, X. Z., and Yan, M. M. H.: A thermal infrared radiation  
1134 parameterization for atmospheric studies, NASA Tech. Memo. NASA/TM-2001-104606, 19,  
1135 68 pp. <https://ntrs.nasa.gov/citations/20010072848>, 2001.
- 1136 Constantinidou, K., Hadjinicolaou, P., Zittis, G., and Lelieveld, J.: Performance of Land Surface  
1137 Schemes in the WRF Model for Climate Simulations over the MENA-CORDEX Domain,  
1138 Earth Systems and Environment, 19, 10.1007/s41748-020-00187-1, 2020.
- 1139 Dee, D. P., Uppala, S. M., Simmons, A. J., Berrisford, P., Poli, P., Kobayashi, S., Andrae, U.,  
1140 Balmaseda, M. A., Balsamo, G., Bauer, P., Bechtold, P., Beljaars, A. C. M., van de Berg, L.,  
1141 Bidlot, J., Bormann, N., Delsol, C., Dragani, R., Fuentes, M., Geer, A. J., Haimberger, L.,  
1142 Healy, S. B., Hersbach, H., Hólm, E. V., Isaksen, L., Kållberg, P., Köhler, M., Matricardi, M.,  
1143 McNally, A. P., Monge-Sanz, B. M., Morcrette, J. J., Park, B. K., Peubey, C., de Rosnay, P.,  
1144 Tavolato, C., Thépaut, J. N., and Vitart, F.: The ERA-Interim reanalysis: configuration and  
1145 performance of the data assimilation system, Quarterly Journal of the Royal Meteorological  
1146 Society, 137, 553-597, 10.1002/qj.828, 2011.

1147 Di Virgilio, G., Evans, J. P., Di Luca, A., Olson, R., Argüeso, D., Kala, J., Andrys, J., Hoffmann, P.,  
1148 Katzfey, J. J., and Rockel, B.: Evaluating reanalysis-driven CORDEX regional climate  
1149 models over Australia: model performance and errors, *Clim. Dyn.*, 53, 2985-3005,  
1150 10.1007/s00382-019-04672-w, 2019.

1151 Di Virgilio, G., Ji, F., Tam, E., Nishant, N., Evans, J. P., Thomas, C., Riley, M. L., Beyer, K., Grose,  
1152 M. R., Narsey, S., and Delage, F.: Selecting CMIP6 GCMs for CORDEX Dynamical  
1153 Downscaling: Model Performance, Independence, and Climate Change Signals, *Earth's*  
1154 *Future*, 10, e2021EF002625, <https://doi.org/10.1029/2021EF002625>, 2022.

1155 Di Virgilio, G., Ji, F., Tam, E., Evans, J. P., Kala, J., Andrys, J., Thomas, C., Choudhury, D., Rocha,  
1156 C., Li, Y., Riley, M.: Evaluation of CORDEX ERA5-forced ‘NARClIM2. 0’ regional climate  
1157 models over Australia using the Weather Research and Forecasting (WRF) model version  
1158 4.1.2, Geoscientific Model Development, <https://doi.org/10.5194/gmd-2024-41>, In review.

1159 DWER.: Climate Adaptation Strategy - Building WA's Climate Resilient Future, Government of  
1160 Western Australia, 25 pages. Online at: [https://www.wa.gov.au/system/files/2023-](https://www.wa.gov.au/system/files/2023-07/climate_adaption_strategy_220623.pdf)  
1161 [07/climate\\_adaption\\_strategy\\_220623.pdf](https://www.wa.gov.au/system/files/2023-07/climate_adaption_strategy_220623.pdf), 2023.

1162 Evans, A., Jones, D., Lellyett, S., and Smalley, R.: An Enhanced Gridded Rainfall Analysis Scheme  
1163 for Australia, Australian Bureau of Meteorology 2020a.

1164 Evans, J. P. and Imran, H. M.: The observation range adjusted method: a novel approach to  
1165 accounting for observation uncertainty in model evaluation, *Environmental Research*  
1166 *Communications*, 6, 071001, 10.1088/2515-7620/ad5ad8, 2024.

1167 Evans, J. P., Ji, F., Lee, C., Smith, P., Argüeso, D., and Fita, L.: Design of a regional climate  
1168 modelling projection ensemble experiment - NARClIM, *Geosci. Model Dev.*, 7, 621-629,  
1169 10.5194/gmd-7-621-2014, 2014.

1170 Evans, J. P., Di Virgilio, G., Hirsch, A. L., Hoffmann, P., Remedio, A. R., Ji, F., Rockel, B., and  
1171 Coppola, E.: The CORDEX-Australasia ensemble: evaluation and future projections, *Clim.*  
1172 *Dyn.*, 10.1007/s00382-020-05459-0, 2020b.

1173 Fiddes, S., Pepler, A., Saunders, K., and Hope, P.: Redefining southern Australia’s climatic regions  
1174 and seasons, *J. South Hemisph. Earth Syst. Sci.*, 71, 92-109, <https://doi.org/10.1071/ES20003>,  
1175 2021.

1176 Giorgi, F.: Thirty Years of Regional Climate Modeling: Where Are We and Where Are We Going  
1177 next?, *Journal of Geophysical Research: Atmospheres*, 124, 5696-5723,  
1178 10.1029/2018jd030094, 2019.

1179 Glotfelty, T., Ramírez-Mejía, D., Bowden, J., Ghilardi, A., and West, J. J.: Limitations of WRF land  
1180 surface models for simulating land use and land cover change in Sub-Saharan Africa and  
1181 development of an improved model (CLM-AF v. 1.0), *Geosci. Model Dev.*, 14, 3215-3249,  
1182 10.5194/gmd-14-3215-2021, 2021.

1183 Grose, M., Narsey, S., Trancoso, R., Mackallah, C., Delage, F., Dowdy, A., Di Virgilio, G.,  
1184 Watterson, I., Dobrohotoff, P., Rashid, H. A., Rauniyar, S., Henley, B., Thatcher, M., Syktus,  
1185 J., Abramowitz, G., Evans, J. P., Su, C.-H., and Takbash, A.: A CMIP6-based multi-model  
1186 downscaling ensemble to underpin climate change services in Australia, *Climate Services*, 30,  
1187 100368, <https://doi.org/10.1016/j.cliser.2023.100368>, 2023.

1188 Grose, M. R., Foster, S., Risbey, J. S., Osbrough, S., and Wilson, L.: Using indices of atmospheric  
1189 circulation to refine southern Australian winter rainfall climate projections, *Clim. Dyn.*,  
1190 10.1007/s00382-019-04880-4, 2019.

1191 Grose, M. R., Narsey, S., Delage, F., Dowdy, A. J., Bador, M., Boschhat, G., Chung, C., Kajtar, J.,  
1192 Rauniyar, S., Freund, M., Lyu, K., Rashid, H. A., Zhang, X., Wales, S., Trenham, C.,  
1193 Holbrook, N. J., Cowan, T., Alexander, L. V., Arblaster, J. M., and Power, S. B.: Insights  
1194 from CMIP6 for Australia's future climate, *Earth's Future*, 8, e2019EF001469,  
1195 <https://doi.org/10.1029/2019EF001469>, 2020.

1196 Herger, N., Abramowitz, G., Knutti, R., Angéilil, O., Lehmann, K., and Sanderson, B. M.: Selecting a  
1197 climate model subset to optimise key ensemble properties, *Earth Syst. Dynam.*, 9, 135-151,  
1198 10.5194/esd-9-135-2018, 2018.

1199 Hong, S.-Y., Noh, Y., and Dudhia, J.: A New Vertical Diffusion Package with an Explicit Treatment  
1200 of Entrainment Processes, *Monthly Weather Review*, 134, 2318-2341,  
1201 <https://doi.org/10.1175/MWR3199.1>, 2006.

1202 Hong, S. Y. and Lim, J.-O. J.: The WRF Single-Moment 6-Class Microphysics Scheme (WSM6),  
1203 *Asia-Pac. J. Atmos. Sci.*, 42, 129-151, 2006.

1204 Hsiang, S., Kopp, R., Jina, A., Rising, J., Delgado, M., Mohan, S., Rasmussen, D. J., Muir-Wood, R.,  
1205 Wilson, P., Oppenheimer, M., Larsen, K., and Houser, T.: Estimating economic damage from  
1206 climate change in the United States, *Science*, 356, 1362-1368, 10.1126/science.aal4369, 2017.

1207 Huang, Y., Xue, M., Hu, X.-M., Martin, E., Novoa, H. M., McPherson, R. A., Perez, A., and Morales,  
1208 I. Y.: Convection-Permitting Simulations of Precipitation over the Peruvian Central Andes:  
1209 Strong Sensitivity to Planetary Boundary Layer Parameterization, *J. Hydrometeorol.*, 24,  
1210 1969-1990, <https://doi.org/10.1175/JHM-D-22-0173.1>, 2023.

1211 Iacono, M. J., Delamere, J. S., Mlawer, E. J., Shephard, M. W., Clough, S. A., and Collins, W. D.:  
1212 Radiative forcing by long-lived greenhouse gases: Calculations with the AER radiative  
1213 transfer models, *Journal of Geophysical Research: Atmospheres*, 113,  
1214 <https://doi.org/10.1029/2008JD009944>, 2008.

1215 Imran, H. M., Kala, J., Ng, A. W. M., and Muthukumaran, S.: An evaluation of the performance of a  
1216 WRF multi-physics ensemble for heatwave events over the city of Melbourne in southeast  
1217 Australia, *Clim. Dyn.*, 50, 2553-2586, 10.1007/s00382-017-3758-y, 2018.

1218 IPCC: Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the  
1219 Sixth Assessment Report of the Intergovernmental Panel on Climate Change, Cambridge  
1220 University Press, 2021.

1221 Iturbide, M., Gutiérrez, J. M., Alves, L. M., Bedia, J., Cerezo-Mota, R., Gimadevilla, E., Cofiño, A.  
1222 S., Di Luca, A., Faria, S. H., Gorodetskaya, I. V., Hauser, M., Herrera, S., Hennessy, K.,  
1223 Hewitt, H. T., Jones, R. G., Krakovska, S., Manzanar, R., Martínez-Castro, D., Narisma, G.  
1224 T., Nurhati, I. S., Pinto, I., Seneviratne, S. I., van den Hurk, B., and Vera, C. S.: An update of  
1225 IPCC climate reference regions for subcontinental analysis of climate model data: definition  
1226 and aggregated datasets, *Earth Syst. Sci. Data*, 12, 2959-2970, 10.5194/essd-12-2959-2020,  
1227 2020.

1228 Janjić, Z. I.: Comments on “Development and Evaluation of a Convection Scheme for Use in Climate  
1229 Models”, *Journal of the Atmospheric Sciences*, 57, 3686-3686, [https://doi.org/10.1175/1520-  
1230 0469\(2000\)057<3686:CODAEO>2.0.CO;2](https://doi.org/10.1175/1520-0469(2000)057<3686:CODAEO>2.0.CO;2), 2000.

1231 Kain, J. S.: The Kain-Fritsch convective parameterization: An update, *Journal of Applied  
1232 Meteorology*, 43, 170-181, 10.1175/1520-0450(2004)043<0170:tkcpau>2.0.co;2, 2004.

1233 Kendon, E. J., Prein, A. F., Senior, C. A., and Stirling, A.: Challenges and outlook for convection-  
1234 permitting climate modelling, *Philosophical transactions. Series A, Mathematical, physical,  
1235 and engineering sciences*, 379, 20190547, 10.1098/rsta.2019.0547, 2021.

1236 Kendon, E. J., Ban, N., Roberts, N. M., Fowler, H. J., Roberts, M. J., Chan, S. C., Evans, J. P., Fosse,  
1237 G., and Wilkinson, J. M.: Do convection-permitting regional climate models improve  
1238 projections of future precipitation change?, *Bulletin of the American Meteorological Society*,  
1239 98, 79-+, 10.1175/bams-d-15-0004.1, 2017.

1240 King, A. D., Alexander, L. V., and Donat, M. G.: The efficacy of using gridded data to examine  
1241 extreme rainfall characteristics: a case study for Australia, *International Journal of  
1242 Climatology*, 33, 2376-2387, 10.1002/joc.3588, 2013.

1243 Kusaka, H. and Kimura, F.: Coupling a Single-Layer Urban Canopy Model with a Simple  
1244 Atmospheric Model: Impact on Urban Heat Island Simulation for an Idealized Case, *Journal  
1245 of the Meteorological Society of Japan. Ser. II*, 82, 67-80, 10.2151/jmsj.82.67, 2004.

1246 Lee, D., Min, S.-K., Ahn, J.-B., Cha, D.-H., Shin, S.-W., Chang, E.-C., Suh, M.-S., Byun, Y.-H., and  
1247 Kim, J.-U.: Uncertainty analysis of future summer monsoon duration and area over East Asia  
1248 using a multi-GCM/multi-RCM ensemble, *Environ. Res. Lett.*, 18, 064026, 10.1088/1748-  
1249 9326/acd208, 2023.

1250 Lucas-Picher, P., Argüeso, D., Brisson, E., Trambly, Y., Berg, P., Lemonsu, A., Kotlarski, S., and  
1251 Caillaud, C.: Convection-permitting modeling with regional climate models: Latest  
1252 developments and next steps, *WIREs Climate Change*, 12, e731,  
1253 <https://doi.org/10.1002/wcc.731>, 2021.



1254 Meehl, G. A., Senior, C. A., Eyring, V., Flato, G., Lamarque, J.-F., Stouffer, R. J., Taylor, K. E., and  
1255 Schlund, M.: Context for interpreting equilibrium climate sensitivity and transient climate  
1256 response from the CMIP6 Earth system models, *Science Advances*, 6, eaba1981,  
1257 10.1126/sciadv.aba1981, 2020.

1258 Murphy, B. F. and Timbal, B.: A review of recent climate variability and climate change in  
1259 southeastern Australia, *International Journal of Climatology*, 28, 859-879,  
1260 <https://doi.org/10.1002/joc.1627>, 2008.

1261 Nakanishi, M. and Niino, H.: Development of an Improved Turbulence Closure Model for the  
1262 Atmospheric Boundary Layer, *Journal of the Meteorological Society of Japan. Ser. II*, 87,  
1263 895-912, 10.2151/jmsj.87.895, 2009.

1264 Nishant, N., Evans, J. P., Di Virgilio, G., Downes, S. M., Ji, F., Cheung, K. K. W., Tam, E., Miller, J.,  
1265 Beyer, K., and Riley, M. L.: Introducing NARCLiM1.5: Evaluating the Performance of  
1266 Regional Climate Projections for Southeast Australia for 1950–2100, *Earth's Future*, 9,  
1267 e2020EF001833, <https://doi.org/10.1029/2020EF001833>, 2021.

1268 Niu, G.-Y., Yang, Z.-L., Mitchell, K. E., Chen, F., Ek, M. B., Barlage, M., Kumar, A., Manning, K.,  
1269 Niyogi, D., Rosero, E., Tewari, M., and Xia, Y.: The community Noah land surface model  
1270 with multiparameterization options (Noah-MP): 1. Model description and evaluation with  
1271 local-scale measurements, *Journal of Geophysical Research: Atmospheres*, 116,  
1272 10.1029/2010jd015139, 2011.

1273 NSW Government.: NSW Climate Change Fund Annual Report 2021-22, 2022.

1274 NSW Government.: NSW Climate Change Fund Annual Report 2022-23, 2023.

1275 Nuryanto, D. E., Satyaningsih, R., Nuraini, T. A., Rizal, J., Heriyanto, E., Linarka, U. A., and  
1276 Sopaheluwakan, A.: Evaluation of Planetary Boundary Layer (PBL) schemes in simulating  
1277 heavy rainfall events over Central Java using high resolution WRF model, *Sixth International  
1278 Symposium on LAPAN-IPB Satellite, SPIE*, 2019.

1279 Oleson, K., Lawrence, D., Bonan, G. B., Flanner, M., Kluzek, E., Lawrence, P., Levis, S., Swenson,  
1280 S. C., Thornton, P. E., Dai, A., Decker, M., Dickinson, R., Feddema, J., Heald, C., Hoffman,  
1281 F., Lamarque, J.-F., Mahowald, N., Niu, G.-Y., Qian, T., and Zeng, X.: Technical Description  
1282 of version 4.0 of the Community Land Model (CLM), 2010.

1283 Pepler, A. and Dowdy, A.: Intense east coast lows and associated rainfall in eastern Australia, *J. South  
1284 Hemisph. Earth Syst. Sci.*, 71, 110-122, 10.1071/es20013, 2021.

1285 Perkins, S. E., Pitman, A. J., Holbrook, N. J., and McAneney, J.: Evaluation of the AR4 climate  
1286 models' simulated daily maximum temperature, minimum temperature, and precipitation over  
1287 Australia using probability density functions, *J. Clim.*, 20, 4356-4376, 10.1175/jcli4253.1,  
1288 2007.

1289 Pleim, J. E.: A Combined Local and Nonlocal Closure Model for the Atmospheric Boundary Layer.  
1290 Part I: Model Description and Testing, *J. Appl. Meteorol. Climatol.*, 46, 1383-1395,  
1291 <https://doi.org/10.1175/JAM2539.1>, 2007.

1292 Rashid, H. A., Sullivan, A., Dix, M., Bi, D., Mackallah, C., Ziehn, T., Dobrohotoff, P., O'Farrell, S.,  
1293 Harman, I. N., Bodman, R., and Marsland, S.: Evaluation of climate variability and change in  
1294 ACCESS historical simulations for CMIP6, *J. South Hemisph. Earth Syst. Sci.*, 72, 73-92,  
1295 <https://doi.org/10.1071/ES21028>, 2022.

1296 Salamanca, F., Zhang, Y. Z., Barlage, M., Chen, F., Mahalov, A., and Miao, S. G.: Evaluation of the  
1297 WRF-Urban Modeling System Coupled to Noah and Noah-MP Land Surface Models Over a  
1298 Semiarid Urban Environment, *Journal of Geophysical Research-Atmospheres*, 123, 2387-  
1299 2408, 10.1002/2018jd028377, 2018.

1300 Sherwood, S. C., Webb, M. J., Annan, J. D., Armour, K. C., Forster, P. M., Hargreaves, J. C., Hegerl,  
1301 G., Klein, S. A., Marvel, K. D., Rohling, E. J., Watanabe, M., Andrews, T., Braconnot, P.,  
1302 Bretherton, C. S., Foster, G. L., Hausfather, Z., von der Heydt, A. S., Knutti, R., Mauritsen,  
1303 T., Norris, J. R., Proistosescu, C., Rugenstein, M., Schmidt, G. A., Tokarska, K. B., and  
1304 Zelinka, M. D.: An Assessment of Earth's Climate Sensitivity Using Multiple Lines of  
1305 Evidence, *Rev. Geophys.*, 58, e2019RG000678, <https://doi.org/10.1029/2019RG000678>,  
1306 2020.

1307 Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Barker, D. M., Wang, W., and Powers, J. G.:  
1308 A description of the Advanced Research WRF Version 3. NCAR Tech Note NCAR/TN-  
1309 475+STR. NCAR, Boulder, CO, 2008.

1310 Tegen, I., Hollrig, P., Chin, M., Fung, I., Jacob, D., and Penner, J.: Contribution of different aerosol  
1311 species to the global aerosol extinction optical thickness: Estimates from model results,  
1312 *Journal of Geophysical Research: Atmospheres*, 102, 23895-23915,  
1313 <https://doi.org/10.1029/97JD01864>, 1997.

1314 Tewari, M., Wang, W., Dudhia, J., LeMone, M. A., Mitchell, K., Ek, M., Gayno, G., Wegiel, J., and  
1315 Cuenca, R.: Implementation and verification of the united NOAH land surface model in the  
1316 WRF model, 11-15 pp.2016.

1317 Thompson, G., Field, P. R., Rasmussen, R. M., and Hall, W. D.: Explicit Forecasts of Winter  
1318 Precipitation Using an Improved Bulk Microphysics Scheme. Part II: Implementation of a  
1319 New Snow Parameterization, *Monthly Weather Review*, 136, 5095-5115,  
1320 <https://doi.org/10.1175/2008MWR2387.1>, 2008.

1321 Tiedtke, M.: A Comprehensive Mass Flux Scheme for Cumulus Parameterization in Large-Scale  
1322 Models, *Monthly Weather Review*, 117, 1779-1800, 10.1175/1520-  
1323 0493(1989)117<1779:acmfsf>2.0.co;2, 1989.

1324 Torma, C., Giorgi, F., and Coppola, E.: Added value of regional climate modeling over areas  
1325 characterized by complex terrain—Precipitation over the Alps, *Journal of Geophysical*  
1326 *Research: Atmospheres*, 120, 3957-3972, 10.1002/2014JD022781, 2015.

1327 WCRP: CORDEX experiment design for dynamical downscaling of CMIP6 (DRAFT),  
1328 [https://cordex.org/wp-content/uploads/2020/06/CORDEX-  
CMIP6\\_exp\\_design\\_draft\\_20200610.pdf](https://cordex.org/wp-content/uploads/2020/06/CORDEX-<br/>1329 CMIP6_exp_design_draft_20200610.pdf), 2020.

1330 WCRP: CORDEX-CMIP6 Data Request, Coordinated Regional Downscaling Experiment  
1331 (CORDEX), [https://cordex.org/wp-content/uploads/2022/03/CORDEX-  
CMIP6\\_Data\\_Request\\_tutorial.pdf](https://cordex.org/wp-content/uploads/2022/03/CORDEX-<br/>1332 CMIP6_Data_Request_tutorial.pdf), 2022.

1333 Whetton, P. and Hennessy, K.: Potential benefits of a “storyline” approach to the provision of regional  
1334 climate projection information, International Climate Change Adaptation Conference,  
1335 NCARF, Gold Coast, Australia, 2010.

1336 Wilks, D. S.: “The Stippling Shows Statistically Significant Grid Points”: How Research Results are  
1337 Routinely Overstated and Overinterpreted, and What to Do about It, *Bulletin of the American*  
1338 *Meteorological Society*, 97, 2263-2273, <https://doi.org/10.1175/BAMS-D-15-00267.1>, 2016.

1339 Xie, K., Li, L., Chen, H., Mayer, S., Dobler, A., Xu, C. Y., and Gokturk, O. M.: Enhanced Evaluation  
1340 of Sub-daily and Daily Extreme Precipitation in Norway from Convection-Permitting Models  
1341 at Regional and Local Scales, *Hydrol. Earth Syst. Sci. Discuss.*, 2024, 1-38, 10.5194/hess-  
1342 2024-68, 2024.

1343 Zhuo, L., Dai, Q., Han, D., Chen, N., and Zhao, B.: Assessment of simulated soil moisture from WRF  
1344 Noah, Noah-MP, and CLM land surface schemes for landslide hazard application, *Hydrol.*  
1345 *Earth Syst. Sci.*, 23, 4199-4218, 10.5194/hess-23-4199-2019, 2019.

1346 Ziehn, T., Chamberlain, M. A., Law, R. M., Lenton, A., Bodman, R. W., Dix, M., Stevens, L., Wang,  
1347 Y.-P., and Srbinovsky, J.: The Australian Earth System Model: ACCESS-ESM1.5, *J. South*  
1348 *Hemisph. Earth Syst. Sci.*, 70, 193-214, <https://doi.org/10.1071/ES19035>, 2020.