

Evaluation of CORDEX ERA5-forced ‘NARClIM2.0’ regional climate models over Australia using the Weather Research and Forecasting (WRF) model version 4.1.2

Giovanni Di Virgilio^{1,2}, Fei Ji^{1,3}, Eugene Tam¹, Jason P. Evans^{2,3}, Jatin Kala⁴, Julia Andrys⁴, Christopher Thomas², Dipayan Choudhury¹, Carlos Rocha¹, Yue Li¹, and Matthew L. Riley¹

¹Climate & Atmospheric Science, NSW Department of Planning and Environment, Sydney, Australia

²Climate Change Research Centre, University of New South Wales, Sydney, Australia

³Australian Research Council Centre of Excellence for Climate Extremes, University of New South Wales, Sydney, Australia

⁴Environmental and Conservation Sciences, and Centre for Climate Impacted Terrestrial Ecosystems, Harry Butler Institute, Murdoch University, Murdoch, WA 6150, Australia

Correspondence to: Giovanni Di Virgilio (giovanni.divirgilio@environment.nsw.gov.au; giovanni@unsw.edu.au)

1 **Abstract.** Understanding regional climate model (RCM) capabilities to simulate current climate
2 informs model development and climate change assessments. This is the first evaluation of the
3 NARClIM2.0 ensemble of [seven](#) Weather Forecasting and Research RCMs driven by ECMWF
4 Reanalysis v5 (ERA5) over Australia at 20 km resolution contributing to CORDEX-CMIP6
5 Australasia, and south-eastern Australia at convection-permitting resolution (4 km). [The performances](#)
6 [of these seven ERA5-RCMs \(R1-R7\)](#), in simulating mean and extreme maximum, minimum
7 temperature and precipitation is evaluated against observations at annual, seasonal, and daily
8 timescales, and compared to corresponding performances of previous-generation CORDEX-CMIP5
9 Australasia ERA-Interim-driven RCMs. ERA5-RCMs substantially reduce cold biases for mean and
10 extreme maximum temperature versus ERA-Interim-RCMs, with [the best-performing ERA5-RCMs](#)
11 [showing](#) small mean absolute biases ([ERA5-R5](#): 0.54K; [ERA5-R1](#): 0.81K, respectively), but produce
12 no improvements for minimum temperature. [At 20 km resolution, improvements in mean and extreme](#)
13 [precipitation for ERA5-RCMs versus ERA-Interim RCMs are principally evident over south-eastern](#)
14 [Australia, whereas strong biases remain over northern Australia. At convection-permitting scale over](#)
15 [south-eastern Australia, mean absolute biases for mean precipitation for the ERA5-RCM ensemble are](#)
16 [around 79% smaller versus the ERA-Interim RCMs that simulate for this region.](#) Although ERA5
17 reanalysis data confer improvements over ERA-Interim, only improvements in precipitation

Deleted: reanalyses

Deleted: performance

Deleted: ERA5-RCM precipitation simulations show lower bias magnitudes versus ERA-Interim-RCMs, though dry biases remain over monsoonal northern Australia and

Deleted: extreme precipitation simulation improvements are

Deleted: evident at convection-permitting 4 km resolution

25 simulation by ERA5-RCMs are attributable to the ERA5 driving data, with RCM improvements for
26 maximum temperature more attributable to model design choices, suggesting improved driving data
27 do not guarantee all RCM performance improvements, with potential implications for CMIP6-forced
28 dynamical downscaling. This evaluation shows that NARClIM2.0 ERA5-RCMs provide valuable
29 reference simulations for upcoming CMIP6-forced downscaling over CORDEX-Australasia and are
30 informative datasets for climate impact studies. Using a subset of these RCMs for simulating CMIP6-
31 forced climate projections over CORDEX-Australasia and/or at convection-permitting scales could
32 yield tangible benefits in simulating regional climate.

Keywords:

33 Climate change; climate impact adaptation; CORDEX-CMIP6; dynamical downscaling; model
34 development; reanalysis

35 1. Introduction

36 Global climate models (GCMs) are optimum tools for simulating future climate at global and
37 continental scales, informing policy and planning at these scales on climate change under different
38 greenhouse gas concentration scenarios (IPCC, 2021). Successive generations of GCMs have seen
39 several improvements, including incremental increases in spatial resolution and some improvements
40 in the simulation of the current climate (Eyring et al., 2016; Stouffer et al., 2017; Grose et al., 2020).
41 However, the coarse spatial resolution of GCMs (100 to 250 km) limits their ability to resolve the
42 fine-scale drivers of regional climate, such as complex topography, land-use, and mesoscale
43 atmospheric processes like convection. This in turn limits their efficacy for climate mitigation and
44 adaptation planning at regional scales (Hsiang et al., 2017).

45 Dynamical downscaling of GCM outputs using regional climate models (RCMs) is one
46 approach for generating high-resolution climate projections at regional scales (Giorgi, 2006; Laprise,
47 2008). RCMs use GCM outputs as initial and lateral boundary conditions to generate fine-scale
48 climate simulations that better resolve the fine-scale drivers of regional climate (Giorgi and Bates,
49 1989; Torma et al., 2015; Di Luca et al., 2012). This can create fine-scale climate information that is
50 spatially and temporally more realistic than the driving GCM information, providing climate
51 simulations more suitable for regional climate impact studies (Giorgi, 2019). However, such
52 improvements are not guaranteed, and typically vary with time and location (Di Virgilio et al., 2019;
53 Di Virgilio et al., 2020b; Panitz et al., 2014; Buchignani et al., 2016). There is also the potential that
54 RCMs simulate climate projections that are not more physically plausible than those of driving GCMs
55 (Ekström et al., 2015). Design considerations such as selection of driving models and RCM
56 parameterisation also underlie the nature of potential improvements in regional climate simulations.

57 The Coordinated Regional Climate Downscaling Experiment (CORDEX) is an initiative of the
58 World Climate Research Programme (WCRP) that provides experimental guidelines facilitating both
59 the production of regional climate projections, and inter-model comparisons across modelling groups
60 (Giorgi et al., 2009). Under CORDEX, regional climate projections based on CMIP5 (Coupled Model
61 Intercomparison Project Phase 5) GCM projections were produced for fourteen regions globally.
62 CORDEX is building on these previous downscaling intercomparison projects to provide a common
63 framework for downscaling activities based on CMIP6 GCMs (Gutowski et al., 2016).

64 A key component of CORDEX is using RCMs to dynamically downscale reanalyses such as
65 ERA-Interim (Dee et al., 2011) under CORDEX-CMIP5, and recently ERA5 (Hersbach et al., 2020)
66 under CORDEX-CMIP6, and evaluating the RCMs' capabilities to simulate present-day climate. If a
67 given RCM performs poorly in simulating the present-day climate, this lowers confidence in future
68 climate changes projected by this model. Assessing the relative strengths and weaknesses of ERA5-
69 forced RCMs can inform the decision to exclude poorer performing RCM configurations when selecting

70 a subset of RCMs for downscaling of CMIP6 GCMs. It also helps benchmark the subsequent
71 performance profiles of CMIP6-forced RCM projections and hindcasts.

72 Previous work to dynamically downscale ERA5 over CORDEX Australasia includes the
73 BARPA-R (Bureau of Meteorology Atmospheric Regional Projections for Australia) regional climate
74 model which simulates over CORDEX Australasia at ~17 km resolution (Howard et al., 2024).
75 Evaluation of BARPA-R's skill in simulating the Australian climate observed good performance
76 overall, including a 1°C cold bias in daily maximum temperatures and wet biases of up to 25 mm/month
77 over inland Australia. Other previous studies of dynamical downscaling of ERA5 by RCMs have
78 focused on short-term (e.g. ~one year) regional climate simulations (e.g. Varga and Breuer, 2020; Zhou
79 et al., 2021) rather than multidecadal simulations. Several have focused on specific regions that are not
80 CORDEX domains, some of which have a smaller spatial extent in comparison. For instance, Reder et
81 al. (2022) conducted dynamical downscaling of ERA5 using COSMO-CLM (CCLM; Rockel et al.
82 2008) on nine separate domains over twenty European cities at convection-permitting scale (~2.2 km).
83 They demonstrated an overall pattern of added value in the simulation of heavy precipitation at city
84 scale relative to the driving reanalysis. Focusing on precipitation simulation over the Lake Victoria
85 Basin in Africa, Van De Walle et al. (2020) conducted ERA5-forced CCLM simulations at convection-
86 permitting scale. They found that CCLM outperformed the ERA5 data set, as well as RCM simulations
87 using parametrised convection, though a domain-averaged wet bias was still evident. These authors
88 attributed the overall improvements in the simulation of sub-daily precipitation to the convection-
89 permitting resolution and improved cloud microphysics. Additionally, two Weather Research and
90 Forecasting model (WRF; Skamarock et al. 2008) experiments over the Tibetan Plateau conducted at
91 'gray-zone' (~9 km) and convection-permitting (~3 km) resolutions for 2009-2018 both showed
92 successful simulation of the spatial pattern and daily variation of surface temperature and precipitation
93 (Ma et al., 2022). Notably, the ability of the convection-permitting WRF RCM in improving
94 precipitation simulation was limited relative to the gray-zone experiment.

95 The sole prior evaluation of reanalysis-driven CORDEX-CMIP5 Australasia regional climate
96 models was conducted by Di Virgilio et al. (2019). This evaluation of CORDEX ERA-Interim forced
97 RCMs focused on four configurations of WRF, and single configurations of CCLM and the
98 Conformal-Cubic Atmospheric Model (CCAM; Mcgregor and Dix, 2008) to simulate the historical
99 Australian climate (1981–2010) at 50 km resolution. These RCMs showed statistically significant,
100 strong cold biases in maximum temperature, which in some cases exceeded -5 K, contrasting with
101 more accurate simulations of minimum temperature, with biases of ± 1.5 K for most WRF
102 configurations and CCAM. The RCMs generally overestimated precipitation, especially over
103 Australia's highly populated eastern seaboard. Notably, Di Virgilio et al. (2019) observed strong
104 negative correlations between simulated mean monthly biases in precipitation and maximum
105 temperature, suggesting that the maximum temperature cold bias was linked to precipitation
106 overestimation.

Deleted: At the time of writing (December 2023), few peer-reviewed studies of dynamical downscaling of ERA5 by RCMs have been published. Many of these

110 This study aims to build on that of Di Virgilio et al. (2019) to present the first evaluation of
111 CORDEX-CMIP6 ERA5-forced WRF RCMs over Australia. It has three main aims: 1) to evaluate the
112 capabilities of seven ERA5-forced WRF RCM configurations to simulate the historical Australian
113 climate, assessing the relative strengths and weaknesses of individual RCMs; 2) compare the
114 performance of current generation CORDEX-CMIP6 ERA5 RCMs with the previous generation of
115 CORDEX-CMIP5 ERA-Interim-forced RCMs following the evaluation approach of Di Virgilio et al.
116 (2019); and 3) investigate whether any performance differences observed for the ERA5-forced
117 relative to the ERA-Interim forced RCMs can be attributed to the change in the driving reanalysis data
118 sets or to other factors, such as the use of different RCM physics configurations and model design
119 specifications. Following Di Virgilio et al. (2019) we evaluate the ability of RCMs to simulate near-
120 surface maximum and minimum air temperature and precipitation at annual, seasonal, and daily time
121 scales. Here, our focus is on evaluating the performances of the different RCM generations, with an
122 investigation of the mechanisms underlying the varying model performances to be the subject of
123 future work.

124 **2. Materials and methods**

125 **2.1 Models**

126 The CORDEX-CMIP5 ERA-Interim forced RCMs (WRF360J, WRF360K, WRF360L, MU-
127 WRFSWWA, CCAM and CCLM) used a domain with quasi-regular grid spacing of approximately 50
128 km ($0.44^\circ \times 0.44^\circ$ on a rotated coordinate system) over the CORDEX-Australasia region. The ERA-
129 Interim WRF RCMs used different versions of WRF: WRF360J-K-L used WRF version 3.6.0,
130 whereas MU-WRFSWWA used version 3.3. ERA-Interim RCM parameterisations for planetary
131 boundary layer physics, surface physics, cumulus physics, land surface model, and radiation, and
132 vertical level settings are shown in Table 1. Three configurations of CORDEX-CMIP5 ERA-Interim
133 WRF RCMs (WRF360J-K-L) were run using two nested domains with one-way nesting. The inner
134 domain located over south-eastern Australia obtained its initial and lateral boundary conditions from
135 an outer domain simulation located over the CORDEX-Australasia region (Figure 1). The inner
136 domain used a resolution of approximately 10 km. Further details on the ERA-Interim-forced RCMs
137 are provided in Di Virgilio et al. (2019), including overviews of the WRF, CCAM and CCLM RCMs.

138 Seven ERA5-forced RCMs comprise the CORDEX-CMIP6 evaluation experiment for
139 NARcliM2.0 (NSW and Australian Regional Climate Modelling), which is the latest generation of
140 NARcliM simulations (Evans et al., 2014; Nishant et al., 2021) [and is one of several RCM ensembles](#)
141 [generating dynamically downscaled climate projections for CORDEX-Australasia \(Grose et al. 2023\)](#).
142 These RCMs were driven by ERA5 boundary conditions for a 42-year period from January 1979 to
143 December 2020. All ERA5 RCMs used WRF version 4.1.2. These CORDEX-CMIP6 ERA5 RCMs

144 were also run using two nested domains with one-way nesting. The outer domain over CORDEX-
145 Australasia used a quasi-regular grid spacing of approximately 20 km (0.2° x 0.2° on a rotated
146 coordinate system), and the inner domain over south-eastern Australia used a resolution of
147 approximately 4 km. Both domains used 45 vertical levels. The seven WRF RCM configurations (R1-
148 R7) used different parameterisations for planetary boundary layer physics, surface physics, cumulus
149 physics, land surface model (LSM), and radiation, noting that several parameters differed relative to
150 those of the ERA-Interim WRF RCMs (Table 1). Four of the ERA5-RCMs used the Noah-MP LSM
151 with its ‘dynamic vegetation cover’ option activated (referred to as ‘dynamic vegetation’ in the WRF
152 users’ guide) (Niu et al., 2011). When deactivated (the default), monthly leaf area index (LAI) is
153 prescribed for various vegetation types and the greenness vegetation fraction (GVF) comes from
154 monthly GVF climatological values. Conversely, when dynamic vegetation cover is activated, LAI
155 and GVF are calculated using a dynamic leaf model. We clarify here that dominant plant-functional
156 types do not change when using this option, but only the LAI and GVF, i.e. only the amount of green
157 cover changes. Additionally, while the indicated cumulus parametrisation was used in the 20 km-
158 resolution outer domain, all ERA5-forced simulations were made convection-permitting in the 4 km
159 inner domain; i.e. no cumulus parametrisation was used. Urban physics was switched on for these
160 simulations. These two design changes are unique to these ERA5-WRF RCMs.

161 The seven ERA5 WRF configurations were selected from an ensemble of seventy-eight
162 structurally different WRF RCMs. Each of these seventy-eight RCMs used different parameterisations
163 for planetary boundary layer, microphysics, cumulus, radiation, and LSM, where parameterisation
164 options were selected via literature review and recommendations from WRF model developers. These
165 seventy-eight test RCMs were run for an entire annual cycle (2016 with a two-month spin-up period
166 commencing 1 November 2015). The seven ERA5 WRF configurations were selected from this larger
167 ensemble based on their skill in simulating the south-eastern Australian climate, whilst retaining as
168 much independent information as possible (Evans et al. 2014; Di Virgilio et al. *in review*).

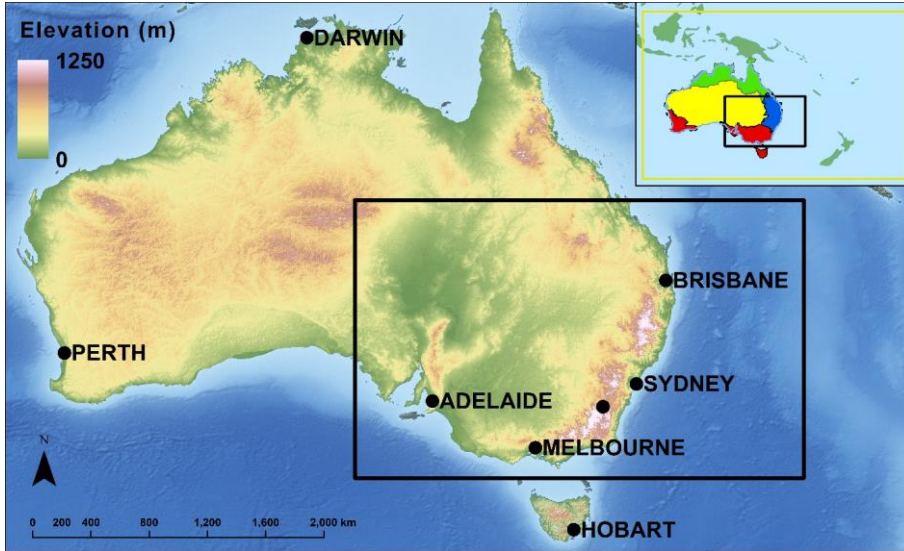
169 Evaluations of model performances are presented for the Australia landmass only and follow the
170 evaluation method of Di Virgilio et al. (2019) for the same period, i.e. for a 29-year period from
171 January 1981 to January 2010. Additionally, select assessments of model performance are presented
172 for the inner domain over south-eastern Australia.

Deleted:

Deleted: W

Deleted: The seven ERA5 WRF configurations were selected from a larger ensemble of seventy-eight WRF RCMs run for an entire annual cycle (2016 with a two-month spin-up period commencing 1 November 2015) based on the criteria that they accurately simulated the south-eastern Australian climate, whilst retaining as much independent information as possible

Deleted: *prep.*



183

184 **Figure 1.** Topographic variation across Australia and major cities Inset: The CORDEX-Australasia
 185 domain and four Natural Resource Management (NRM) regions/climate zones (blue = Eastern
 186 Australia; red = Southern Australia; yellow = Rangelands; and green = Northern Australia). Seven
 187 configurations of CORDEX-CMIP6 ERA5 weather research and forecasting (WRF) RCMs (R1-R7)
 188 and three configurations of CORDEX-CMIP5 ERA-Interim WRF RCMs (WRF360J-K-L) were run
 189 using two nested domains via one-way nesting with an outer domain over CORDEX Australasia and
 190 an inner domain over south-eastern Australia (black rectangle in both main panel and inset).

191 **Table 1.** List of CORDEX-CMIP6 ERA5 and CORDEX-CMIP5 ERA-Interim forced RCMs assessed
 192 by this evaluation study.

Reanalysis	RCM / Version	Planetary boundary layer physics / surface layer physics	Microphysics	Cumulus physics	Shortwave and longwave radiation physics	Land surface	Land options	Vertical Levels
ERA5	R1	YSU (Hong et al., 2006)	WSM6 (Hong and Lim, 2006)	BMJ (Janjić, 2000)	New Goddard (Chou et al., 2001)	Noah Unified (Tewari et al., 2016)	N/A	45
	R2	MYNN2 (Nakanishi and Niino, 2009)	WSM6	Kain-Fritsch (Kain, 2004)	RRTMG (Iacono et al., 2008)	Noah-MP (Niu et al., 2011)	dynamic vegetation	
	R3	MYNN2	Thompson (Thompson et al., 2008)	BMJ	RRTMG	Noah-MP	dynamic vegetation	
	R4	MYNN2	Thompson	BMJ	RRTMG	Noah-MP	TOPMODEL runoff (SIMGM groundwater)	
	R5	ACM2 (Pleim, 2007)	Thompson	BMJ	RRTMG	Noah-MP	dynamic vegetation	
	R6	ACM2	Thompson	Tiedtke (Tiedtke, 1989)	RRTMG	Noah-MP	dynamic vegetation	
	R7	ACM2	Thompson	Tiedtke	RRTMG	Noah-MP	TOPMODEL runoff (SIMGM groundwater)	

	WRF360J	Mellor-Yamada-Janjic/ETA Similarity	WRF Double-Moment 5	Kain-Fritsch	Dudhia/RRTM	Noah Unified	
	WRF360K	Mellor-Yamada-Janjic/ETA Similarity	WRF Double-Moment 5	Betts-Miller-Janjic	Dudhia/RRTM	Noah Unified	
	WRF360L	Yonsei University/MM5 Similarity	WRF Double-Moment 5	Kain-Fritsch	CAM3/CAM3	Noah Unified	30
ERA-I	SWWA WRF330	Yonsei University/MM5 Similarity	WRF Single-Moment 5	Kain-Fritsch	Dudhia/RRTM	Noah Unified	N/A
	CCAM	Monin-Obukhov Similarity Theory stability-dependent boundary-layer scheme (McGregor 1993)	Liquid and ice-water scheme (Rotstayn 1997)	Mass-flux closure (McGregor 2003)	GFDL (Freidenreich and Ramaswamy 1999)	CABLE (Kowalczyk et al. 2006)	27
	CCLM4-8-17-CLM3-5	Prognostic turbulent kinetic energy (Raschendorfer 2001)	Seifert and Beheng (2001), reduced to one moment scheme	Bechtold et al. (2008)	Ritter and Geleyn (1992)	CLM; (Dickinson et al. 2006)	35

193

194 2.2 Observations

195 Australian Gridded Climate Data (AGCD version 1.0; Bureau of Meteorology, 2020; Evans et al., 2020)
196 were used to evaluate RCM performance. This daily gridded maximum and minimum temperature and
197 precipitation data set has a grid-averaged resolution of 0.05° and is obtained from an interpolation of
198 station observations across the Australian continent. Observations include temperature minima and
199 maxima only; hence, the ability of RCMs to reproduce mean temperature was not assessed. Following
200 Di Virgilio et al. (2019), the AGCD data were re-gridded to correspond with the RCM data on their
201 native grids using a conservative area-weighted re-gridding scheme. Most stations used for AGCD are
202 in coastal areas, contrasting with a sparser representation inland, and especially in Australia's north-
203 west. There are more precipitation stations than temperature stations. Only land points over Australia
204 were evaluated because AGCD observations are terrestrial data.

205 2.3 Evaluation methods

206 2.3.1 Evaluations of CORDEX-CMIP6 ERA5 RCMs versus CORDEX-CMIP5 ERA- 207 Interim RCMs

208 Annual and seasonal means were calculated for maximum and minimum temperature and precipitation
209 using monthly averages for each temperature variable, and the monthly sum for precipitation.
210 Percentiles (i.e. extremes: 99th percentiles for maximum temperature and precipitation; 1st percentile for
211 minimum temperature) were calculated using daily values. RCM performances in reproducing
212 observations over these timescales were assessed by calculating the model bias, i.e. model outputs
213 minus observations, and the RMSE of modelled versus observed fields. The statistical significance of
214 mean annual and seasonal biases compared to the AGCD observations was calculated for each grid cell
215 using t-tests ($\alpha = 0.05$) for maximum and minimum temperature assuming equal variance. The Mann-
216 Whitney U test was used for precipitation given its non-normality. Results on the statistical significance

217 of each ensemble mean were separated into three categories following Tebaldi et al. (2011): 1)
218 statistically insignificant areas are shown in colour, denoting that less than 50% of RCMs are
219 significantly biased, which is the most desired outcome; 2) in areas of significant agreement (stippled),
220 at least 50% of RCMs are significantly biased and at least 66% of significant models agree on the sign
221 of the bias. In such areas, many ensemble members have the same bias sign which is an undesirable
222 outcome; and 3) areas of significant disagreement are shown in white, where at least 50% of RCMs are
223 significantly biased and fewer than 66% of significant models agree on the bias sign.

224 The ability of the RCMs to simulate observed variables at daily time scales was also assessed
225 by comparing the probability density functions (PDFs) for daily mean observations versus those of the
226 RCMs. PDFs were separately calculated for Australia and for each of four natural resource management
227 (NRM) climate regions shown in Figure 1 for maximum and minimum temperature, and precipitation.
228 Here, daily precipitation values below 0.1 mm were omitted from the RCM output, because rates below
229 this amount fall below the detection limit of the stations used to produce the observed data set.
230 Additionally, the daily rainfall observational network used to produce the AGCD has large gaps in
231 several areas of central Australia; hence, RCM output was masked over these areas. RCM and observed
232 PDFs were compared using the Perkins Skill Score (PSS; Perkins et al. (2007), which measures the
233 degree of overlap between two PDFs, with PSS = 1 indicating that the distributions overlap perfectly.

234 **2.3.2 Comparing ERA5 versus ERA-Interim RCM performances after switching driving** 235 **reanalyses**

236 Any performance differences of the ERA5-forced and ERA-Interim-forced RCMs could be partially
237 due to the change in the driving reanalysis, as well as factors such as different RCM physics
238 configurations, model version and other design specifications. To assess whether the change in ERA5
239 versus ERA-Interim driving reanalyses may underlie differences in performance profiles of the WRF
240 RCMs from the two generations of CORDEX experiment we conduct two investigations: 1) the ERA5
241 and ERA-Interim reanalysis data are compared against AGCD observations to assess their degree of
242 bias for annual and seasonal timescales; and 2) fourteen-month simulations are performed where
243 otherwise identically parameterised and configured CORDEX-CMIP6 NARClM2.0 R1-R7 RCMs are
244 forced by ERA-Interim as opposed to ERA5, and similarly the WRFJ-K-L RCMs from the CORDEX-
245 CMIP5 era are forced with ERA5 instead of ERA-Interim. For instance, the ERA5-RCMs CORDEX-
246 CMIP6 (NARClM2.0) RCMs are run for the same 4 km convection permitting domain using the same
247 physics options and model setups with the only changes being to swap ERA5 for ERA-Interim and
248 running for 14 months. These simulations start on 1 November 2015, with evaluation performed for the
249 twelve months of 2016, i.e. using the first 2-months as spin-up period. Australia experienced a range of
250 weather extremes during 2016 driven by a range of climatic influences making 2016 a suitable target
251 year (Bureau of Meteorology, 2017). Owing to finite compute resources, it was not possible to simulate
252 for a longer period for these experiments.

Deleted: for the whole domain for maximum and minimum temperature, and precipitation

Deleted: the

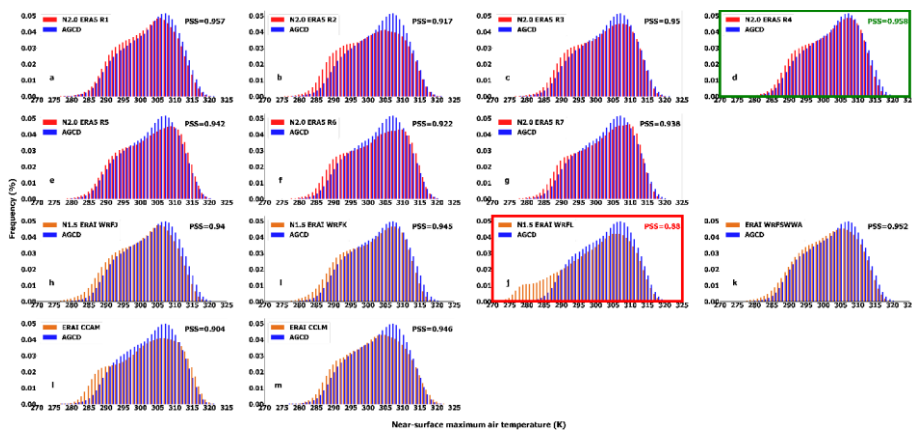
256 **3. Results**

257 RCM evaluation results are presented first for the 29-year CORDEX-CMIP6 ERA5-forced and
 258 CORDEX-CMIP5 ERA-Interim-forced simulations. Evaluation results from switching the driving
 259 reanalyses of the CORDEX-CMIP6 and CORDEX-CMIP5 RCMs are then considered.

260 **3.1 Evaluation of CORDEX-CMIP6 ERA5-RCM and CORDEX-CMIP5**
 261 **ERA-Interim performances**

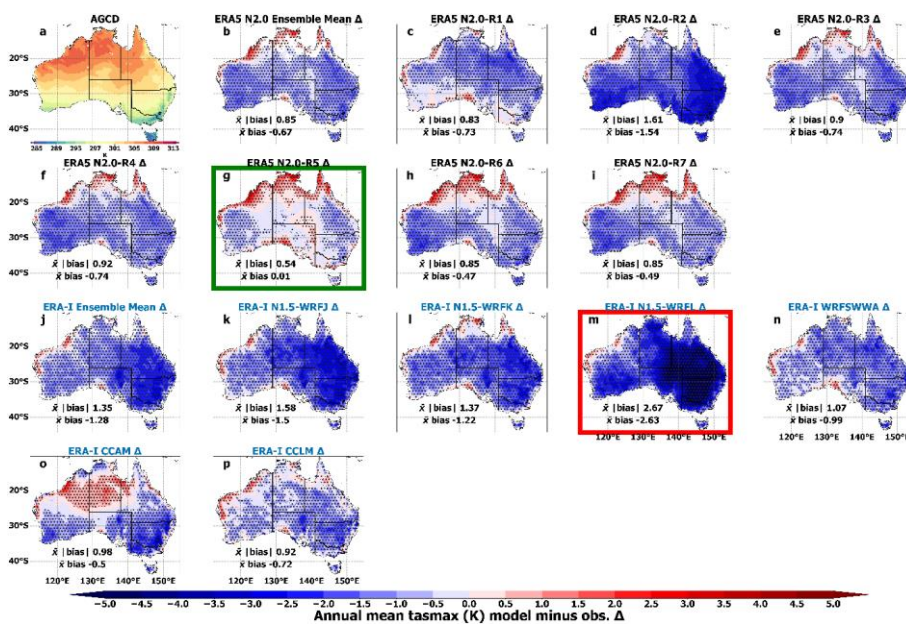
262 **3.1.1 Maximum Temperature**

263 Both ERA5 and ERA-Interim forced RCMs overestimate the frequency of lower-than-average
 264 maximum temperatures and underestimate the observed peaks (Fig. 2). However, most ERA5 RCMs
 265 simulate occurrences of warmer than average temperatures more accurately than the ERA-Interim
 266 RCMs, especially ERA5-R3 (Fig. 2c). The ERA5-RCMs with highest PSS scores (i.e. >0.95; R1 and
 267 R4) show closer correspondences to the observed peaks than the other ERA5-RCMs, but they
 268 underestimate the distribution right tail. [In some respects, RCM performances in PDFs stratified by](#)
 269 [NRM region can show different patterns of results versus the nationally aggregated data \(Online](#)
 270 [Resource 1: Figures S1-S4\). For instance, most ERA5-RCMs show larger over-estimations of warmer](#)
 271 [than average daily maximum temperatures over the Northern Australia region \(Figure S4\) than for](#)
 272 [Australia-wide data \(Figure 2\).](#)



273
 274 **Figure 2.** Probability density functions (PDFs) of mean daily maximum near-surface air temperatures
 275 (K) across Australia for 1981-2010. Panels a-m show the PDF of a specific RCM configuration
 276 relative to that of Australian Gridded Climate Data (AGCD) observations; a-g are NARClm2.0
 277 ERA5-forced RCM configurations; h-m are ERA-Interim-forced RCM configurations. Panel
 278 boundaries in green (red) indicate the RCMs with highest (lowest) PSS. [PDF bin width is 1 K.](#)

279 Most ERA5-RCMs show small cold biases of ~ 0.5 to 1 K for annual mean maximum
 280 temperature over most of Australia, except for warm biases of ~ 0.5 to 1.5 K over the coastal north,
 281 depending on location/RCM configuration (Fig. 3 b-i). ERA5-R5-R7 show lowest area-averaged
 282 absolute annual biases, with R5 showing very small biases of < 0.5 K over much of eastern Australia
 283 (Fig. 3g). ERA5-R2 shows markedly poorer performance than every other ERA5 RCM, with cold
 284 biases exceeding 2 K in some areas (Fig. 3d). The positive biases of maximum temperature over the
 285 tropics for several of the ERA5-RCMs generally correspond well to negative precipitation biases over
 286 this region (see Fig. 7b; e-i). Except for ERA5-R2, the ERA5-forced RCMs show considerable
 287 reductions in the magnitude of cold bias relative to the ERA-Interim forced RCMs (Fig.3 j-p). The
 288 best-performing ERA5-RCM (R5) has an area-averaged absolute mean bias of 0.54 K, as compared to
 289 0.92 K for the best performing ERA-Interim RCM (CCLM), a 52% percentage difference. ERA5-R5
 290 has a 66% percentage difference in absolute bias compared to the best performing ERA-Interim WRF
 291 RCM (i.e. WRFSSWA: 1.07 K).



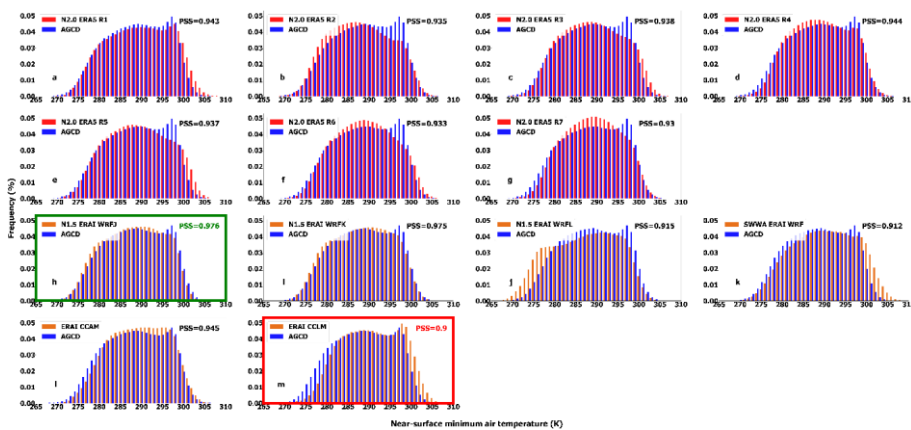
292
 293 **Figure 3.** Annual mean near-surface atmospheric maximum temperature bias with respect to
 294 Australian Gridded Climate Data (AGCD) observations for 1981-2010. Stippled areas indicate
 295 locations where an RCM shows statistically significant bias ($P < 0.05$). b Significance stippling for
 296 the ensemble mean bias follows Tebaldi et al. (2011) and is applied separately to each of the two
 297 RCM ensembles. Statistically insignificant areas are shown in colour, denoting that less than half of
 298 the models are significantly biased. In significant agreeing areas (stippled), at least half of RCMs are
 299 significantly biased, and at least 66% of significant RCMs in each ensemble agree on the direction of
 300 the bias. Significant disagreeing areas are shown in white, which are where at least half of the models
 301 are significantly biased and less than 66% of significant models in each ensemble agree on the bias
 302 direction - see main text for additional detail on the stippling regime. Panel boundaries in green (red)
 303 indicate the RCMs with lowest (highest) area-averaged mean absolute biases

304 During summer, the magnitude and spatial extent of maximum temperature warm biases
 305 increase for all RCMs relative to the annual mean biases (Fig. S5). During winter, several ERA5
 306 RCMs (R1, R3, R4, R5) retain much smaller cold biases than most ERA-Interim-forced models (Fig.
 307 S6). RMSE magnitudes peak for most ERA5 and ERA-Interim models in February (at the end of
 308 austral summer), except for several ERA-Interim RCMs which show larger RMSEs in winter,
 309 especially ERAI-WRFL; Fig. S7).

310 For extreme (99th percentile) maximum temperatures, whilst ERA5-RCMs show lower overall
 311 biases relative to the ERA-Interim RCMs, the former show strong warm biases along coastlines that
 312 are typically stronger than biases further inland (Fig. S8). These biases are particularly pronounced
 313 along northern and eastern coastlines. ERA5-R1 and R5 show the lowest overall mean absolute biases
 314 for extreme maximum temperature, especially over south-eastern Australia. [The various mean
 315 absolute bias and PSS statistics for maximum temperature for the 20 km domain are summarised in
 316 Online Resource Table S1.](#)

317 3.1.2 Minimum Temperature

318 PDFs of daily minimum temperature for the ERA-Interim-forced WRFJ and WRFK RCMs match
 319 observations most closely relative to the ERA5- and other ERA-Interim forced RCMs (Fig. 4).
 320 Observed PDFs [at the continental scale](#) show a slight bimodality that is captured by ERA5-R1, ERA5-
 321 R4, ERAI-WFJ, ERAI-SWWA and ERAI-CCLM. [However, this bimodality is generally not present
 322 in PDFs stratified for specific NRM regions \(Figures S9-S12\).](#) Several RCMs struggle to simulate
 323 minimum temperature occurrences in the middle of the distribution (i.e. ~285-290K), except for
 324 ERA5-R5 and ERA-Interim-WRFJ, WRFK, and CCLM which closely match minimum temperatures
 325 in this range.



326
 327 **Figure 4.** Probability density functions (PDFs) of mean daily minimum near-surface air temperatures
 328 (K) across Australia for 1981-2010. Panels a-m show the PDF of a specific RCM configuration
 329 relative to that of Australian Gridded Climate Data (AGCD) observations; a-g are NARCIIM2.0

Deleted: Online Resource 1:

Deleted: 1

Deleted: 2

Deleted: 3

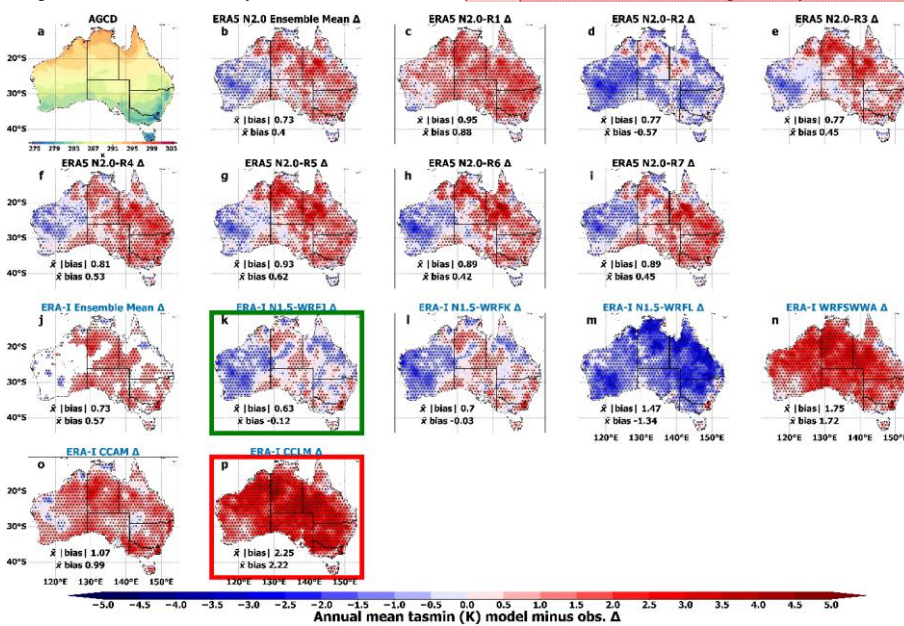
Deleted: (99th percentile)

Deleted: 4

Deleted: only

337 ERA5-forced RCM configurations; h-m are ERA-Interim-forced RCM configurations. Panel
 338 boundary colouring as per Fig. 2. PDF bin width is 1 K.

339 ERA5-RCMs generally overestimate mean minimum temperature annually (Fig. 5) and
 340 seasonally (Fig S13-summer and S14-winter), except for ERA5-R2 which is cold biased. In contrast,
 341 ERA-Interim-RCMs show a mixed signal for WRF-J and WRF-K, cold bias for WRF-L and warm
 342 biases for the remaining RCMs. Warm biases are strongest during JJA for most ERA5-RCMs, and
 343 especially for ERA-Interim CCAM and CCLM (Fig. S14). Whereas ERA5-R2 performs generally
 344 poorly for maximum temperature relative to the other ERA5-RCMs (e.g. annual mean absolute bias =
 345 1.61K), its bias is substantially reduced for minimum temperature (annual mean absolute bias =
 346 0.77K). ERA5 R2 and R3 show better performance for minimum temperature relative to the other
 347 ERA5-RCMs. Their area-averaged annual mean absolute biases (0.77K in both cases) are more
 348 comparable to the ERA-Interim-forced WRFJ-K RCMs which simulate annual mean minimum
 349 temperature most accurately (annual mean absolute biases = 0.66K and 0.7 K, respectively).



350
 351 **Figure 5.** Annual mean near-surface atmospheric minimum temperature bias with respect to gridded
 352 observations for 1981-2010. Stippling and panel boundary colouring as per Fig. 3

353 RMSE annual cycles for mean minimum temperature broadly reflect the above pattern of
 354 results (Fig. S15). For most months throughout the annual cycle, RMSEs are typically lowest for
 355 ERA-Interim WRFJ-K. However, ERA5-R1, R2 also show small RMSEs from May to August, with
 356 RMSEs also being low for ERA5-R3 during spring (September to November).

Deleted: 5

Deleted: 6

Deleted: 6

Deleted: |

Deleted: |

Deleted: |

Deleted: |

Deleted: |

Deleted: |

Deleted: |

Deleted: |

Deleted: S7

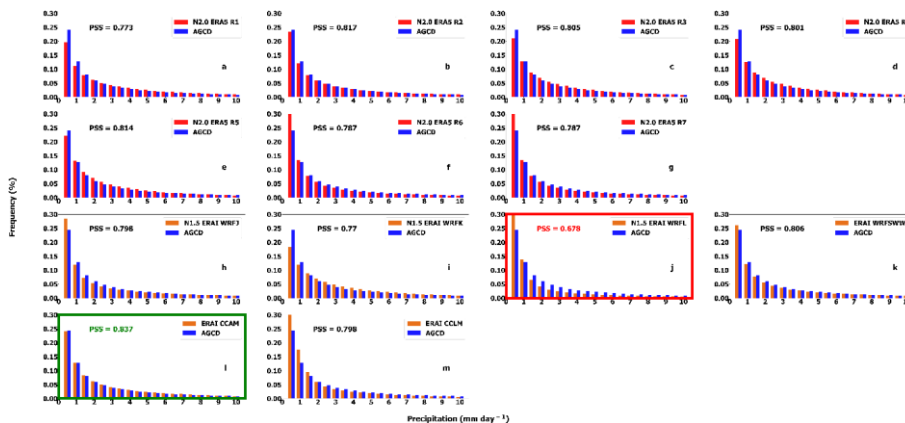
369 The majority of ERA5 and ERA-Interim RCMs are generally warm-biased for extreme
 370 minimum temperature over most of Australia, with only small areas of cold bias over the north-west
 371 (Fig. S16). The exceptions are ERA5-R2 and ERA-Interim-WRFJ-K which show biases of mixed sign
 372 across larger areas of Australia, and ERA-Interim WRFL which is strongly cold biased (Fig. S16).
 373 ERA5-R2 and R3 show reasonably good performance for extreme minimum temperature as compared
 374 to the other ERA5 models, however, ERA-Interim WRFJ-K simulate extreme minimum temperature
 375 most accurately. Mean absolute bias and PSS statistics for minimum temperature for the 20 km
 376 domain are summarised in Table S1.

Deleted: 8

Deleted: 8

377 3.1.3 Precipitation

378 PDFs of mean daily precipitation show that ERA5-R2, ERA-Interim-forced CCAM and WRFSWWA
 379 simulate the occurrence of rainfall events up to 5 mm day⁻¹ more accurately than the other RCMs (Fig.
 380 6). Heavier rainfall events (approximately >7 mm day⁻¹) are underestimated by several RCMs.
 381 Overall, the ERA5-RCMs simulate daily precipitation occurrences consistently better than the ERA-
 382 Interim-RCMs, i.e. four of the seven ERA5-RCMs have PSS >0.8 compared to two of six ERA-
 383 Interim RCMs. Of the ERA5-forced RCMs, R2 produces the best simulation of daily rainfall
 384 occurrences. There are some interesting differences in RCM performance between the NRM regions
 385 (Fig. S17-S20). For instance, most RCMs generally show more skill in capturing daily precipitation
 386 distributions over Southern Australia than other NRM regions, with the ERA5-RCMs performing
 387 particularly well over this region (Fig. S18). Conversely, performances of most RCMs are generally
 388 poorer over Northern Australia than other regions, though ERA5-R5 and ERA-Interim-CCAM show
 389 better performances than their peers over this region with PSS of 0.743 and 0.746, respectively, versus
 390 mean PSS of 0.697 (standard deviation = 0.058; Fig. S20).

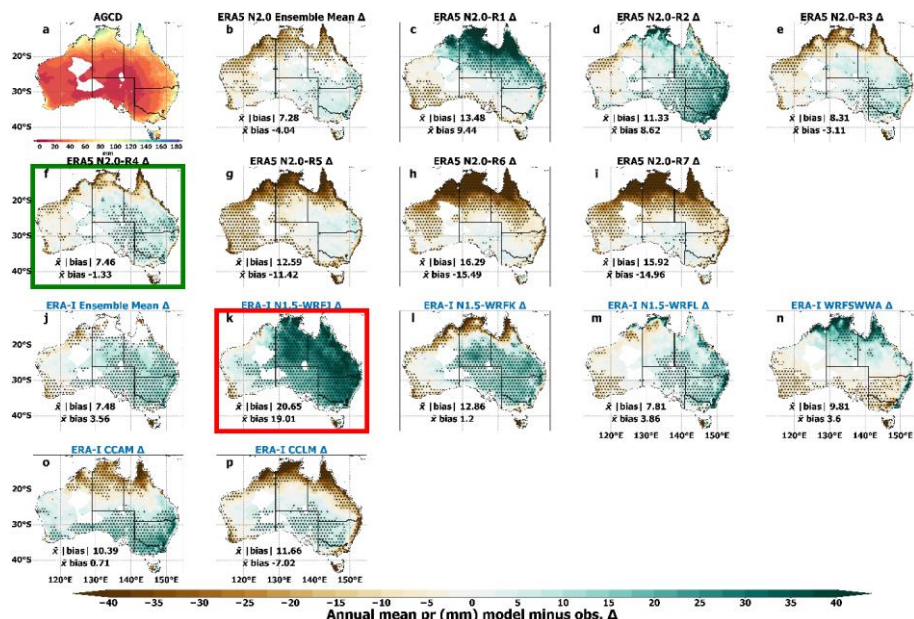


391
 392 **Figure 6.** Probability density functions (PDFs) of mean daily precipitation (mm day⁻¹) across
 393 Australia for 1981-2010. Panels a-m show the PDF of a specific RCM configuration relative to that of
 394 Australian Gridded Climate Data (AGCD) observations; a-g are NARcliM2.0 ERA5-forced RCM

397 configurations; h-m are ERA-Interim-forced RCM configurations. Panel boundary colouring as per
 398 Fig. 2. PDF bin width is 0.5 mm.

399 All ERA5 RCMs except for R1 and R2 are dry-biased for annual mean precipitation over the
 400 monsoonal north (Fig. 7), with R6-7 producing the strongest dry biases exceeding -40 mm over this
 401 region (Fig. 7h-i). Of the ERA5 RCMs, R1 and R2 are exceptional in that they show widespread wet
 402 biases. ERA5-R1 and R2 both use WSM6 microphysics, whereas R3-R7 use Thompson microphysics
 403 (see Discussion 4.1). ERA5-R2 shows the strongest wet-bias over eastern Australia of ~20 mm,
 404 whereas ERA5-R3-4 show smaller wet biases (~5-10 mm) over this region. All ERA5-forced models
 405 show dry biases (between -20 and -35 mm) along the south-western coastline of western Australia.
 406 Overall, with the exceptions of R6 and R7, the ERA5-forced RCMs show reduced mean precipitation
 407 bias relative to the ERA-Interim forced RCMs, especially over southeastern Australia. All RCMs
 408 show the strongest biases (of either sign) during DJF (Fig. S21). For instance, the area and magnitude
 409 of dry-bias over northern Australia increase for ERA5-R3-R7 (Fig. S21). All RCMs show the smallest
 410 biases during JJA (Fig. S22).

Deleted: S9
 Deleted: both
 Deleted: 9
 Deleted: 10



411
 412 **Figure 7.** Annual mean precipitation bias with respect to gridded observations for the RCMs for
 413 1981-2010. Stippling and panel boundary colouring as per Fig. 3.

414 Overall, RMSE annual cycles are similar for the different RCMs (Fig. S23). ERA-Interim
 415 CCAM has the lowest RMSEs throughout the year. Otherwise, all ERA5-forced RCMs have lower
 416 RMSEs than the ERA-Interim forced models (except for CCAM) from April to October, which is an
 417 important growing season in southern Australia.

Deleted: 11

423 The ERA5-RCMs generally over-estimate extreme precipitation over Australia and especially
424 the south-east, though R3, R4 and R5 show widespread dry biases over north-western regions (Fig.
425 S24). The R1 and R2 RCMs show larger extreme precipitation wet biases relative to the other ERA5
426 RCMs (i.e. mean absolute biases of 20.02 mm and 14.83 mm, versus 9.21 mm to 11.4 mm, Fig. S24).
427 Several ERA-Interim-forced RCMs (i.e. WRFJ, WRFK, WRFL) produce similar patterns of bias to
428 the ERA5 RCMs, for instance, with wet biases over south-eastern Australia and dry biases over
429 northern and central regions. Overall, the magnitude of biases over the outer domains is similar
430 between the different RCM generations, with several RCMs showing low mean absolute biases,
431 ranging from 8.75 mm to 10.25 mm. However, focusing specifically on the high-resolution inner
432 domains of ERA5-RCMs and ERA-Interim-WRFJ-K-L RCMs, noting this domain is uniquely
433 convection-permitting (~4 km) for ERA5-RCMs, most ERA5-RCMs show smaller biases than WRFJ-
434 K-L (Fig. S25). For this inner domain, ERA5-R3, R5, R6, R7 show small biases (i.e. <9 mm),
435 particularly over south-eastern coastal areas. Mean absolute bias and PSS statistics for precipitation
436 for the 20 km domain are summarised in Table S1.

437 3.2 Assessing the effects of switching driving ERA5 versus ERA-Interim 438 reanalyses on RCM performances

439 This section investigates whether performance differences of the ERA5-forced and ERA-Interim-
440 forced RCMs may be attributable to the different generations of driving reanalyses as opposed to
441 factors such as different RCM physics parameterisations and design specifications. First, biases in the
442 two reanalyses data sets with respect to observations are assessed. The assessment then focuses on the
443 capacities of the CORDEX-CMIP6 era R1-R7 RCMs and the CORDEX-CMIP5 era WRFJ-K-L
444 RCMs to simulate the south-eastern Australian climate when each RCM generation uses first ERA5
445 and then ERA-Interim driving data. This assessment also provides a further view of the how the WRF
446 RCM performances vary over this high-resolution domain relative to the CORDEX Australasia
447 domain. These comparative simulations are only available for the higher resolution inner domain over
448 south-eastern Australia.

449 3.2.1 ERA5 and ERA-Interim reanalysis biases relative to observations

450 Both ERA5 and ERA-Interim are generally cold biased in their simulation of mean maximum
451 temperature at annual, summer and winter timescales during 1981-2010 (Fig. S26). However, biases
452 are larger in magnitude for ERA-Interim relative to ERA5, especially during summer i.e. ERA5 mean
453 absolute bias = 1.22 K; ERA-Interim = 2.07 K. Biases in ERA5 and ERA-Interim during 2016 are
454 largely consistent with these results (Fig. S27).

455 ERA5 and ERA-Interim overestimate mean minimum temperature over most of Australia at
456 all timescales for both 1981-2010 (Fig. S28) and 2016 (Fig. S29). Biases are again smaller for ERA5

Deleted: 1

Deleted: |

Deleted: |

Deleted: 1

Deleted: |

Deleted: |

Deleted: WRF

Deleted: 1

Deleted: very

Deleted: 5

Deleted: 14

Deleted: |

Deleted: |

Deleted: 15

Deleted: 16

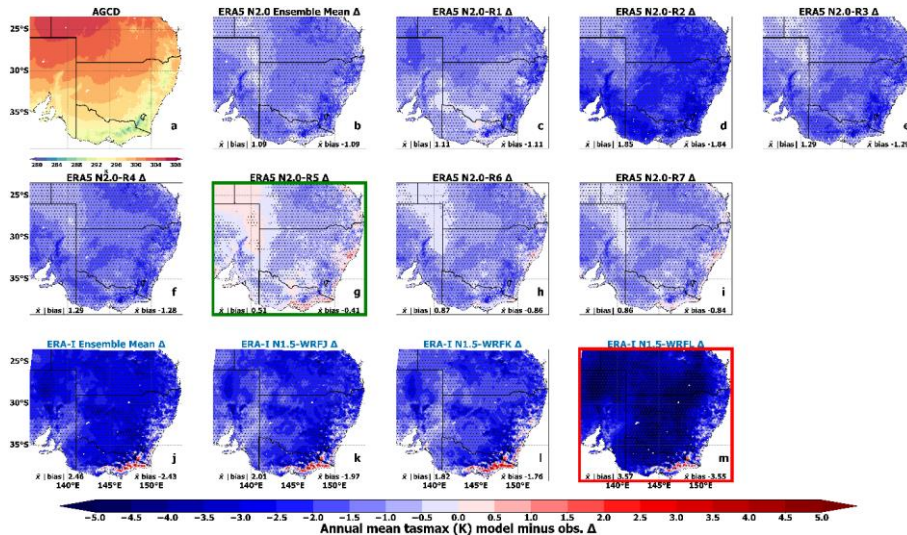
Deleted: 17

473 than for ERA-Interim. For ERA-Interim, warm biases are especially large in magnitude along the
 474 eastern and southern coastlines and over the island of Tasmania.

475 ERA5 shows substantial improvements in simulating mean precipitation at all timescales
 476 relative to ERA-Interim (Fig. S30, i.e. ERA5 annual mean absolute bias = 4.18 mm; ERA-Interim =
 477 8.14 mm). This applies to both periods assessed, i.e. including for 2016 (Fig. S31). Additional
 478 differences in the biases between the reanalysis data sets include ERA-Interim's stronger dry biases
 479 over the monsoonal north during summer (wet season) and marked dry biases along the eastern
 480 coastline and elevated terrain in south-eastern Australia (Fig. S30).

481 3.2.2 Comparing RCM performances after switching the driving reanalyses

482 Prior to switching the driving reanalyses of the two generations of RCMs, the ERA5-NARClIM2.0
 483 RCMs show large reductions in cold bias (Fig. 8b-i) relative to the ERA-Interim-forced RCMs (Fig.
 484 8j-m), with ensemble mean bias magnitudes of 1.09K and 2.46K, respectively.



485 **Figure 8.** Annual mean near-surface atmospheric maximum temperature bias simulated over south-
 486 eastern Australia (WRF simulation inner domain) with respect to gridded observations for the period
 487 1981-2010 for NARClIM2.0 RCMs (b-i) and NARClIM1.5 RCMs (j-m). Stippling and panel
 488 boundary colouring as per Fig. 3.

490 Switching the driving reanalysis of the CORDEX-CMIP6 NARClIM2.0 RCMs, shows small
 491 improvements in the simulation of maximum temperature for several ERA-Interim-forced
 492 NARClIM2.0 RCMs (i.e. for R1, R2, R3 and R7; Fig. 9c,d,e,i). In contrast, ERA-Interim-
 493 NARClIM2.0 R4-5-6 show slight degradations in performance (Fig. 9f,g,h). However, the
 494 NARClIM2.0 ERA-Interim ensemble mean average absolute bias is 0.91K versus 1.09K for the
 495 NARClIM2.0 ERA5 ensemble. Therefore, overall, there is a small performance improvement in

Deleted: 18

Deleted: |

Deleted: |

Deleted: 19

Deleted: 18

Deleted: Without

Deleted: ERA5-forced CORDEX-CMIP6 'NARClIM2.0' RCMs and ERA-Interim CORDEX-CMIP5 RCMs simulate annual mean maximum temperature over the inner domains (Fig. 8) in a similar manner as compared to over Australia (Fig. 3). That is,

Deleted: in the marked

Deleted: es

Deleted: that characterise

Deleted: |

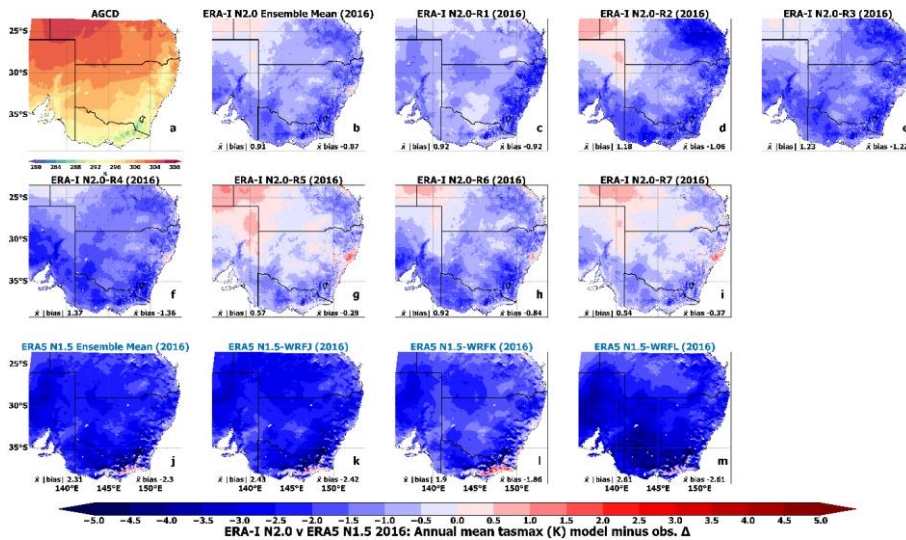
Deleted: es|

Deleted: generation

Deleted: |

Deleted: |

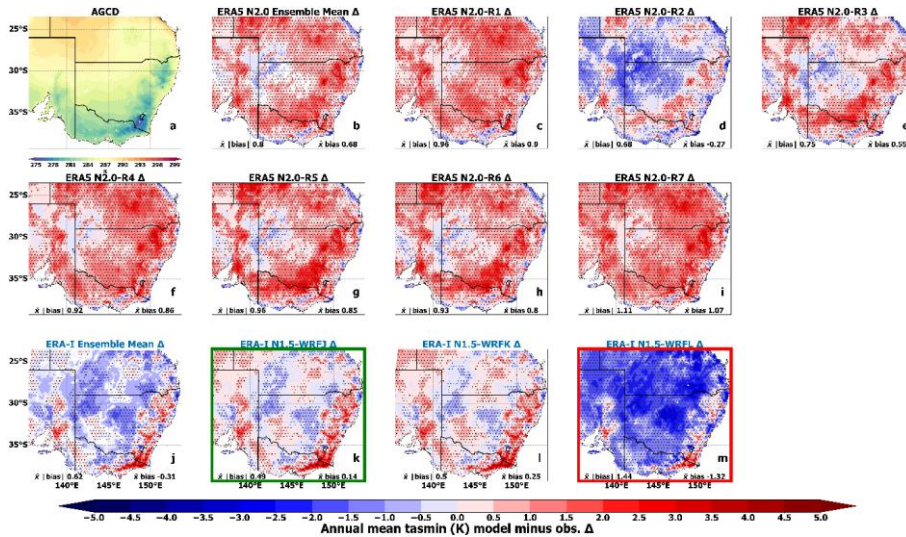
515 forcing the CORDEX-CMIP6 era RCMs using the older reanalysis. Similarly, the CORDEX-CMIP5
 516 era WRFJ and WRFK show poorer simulations of maximum temperature when forced using ERA5
 517 (Fig. 9k-l) relative to their ERA-Interim-forced counterparts, with only ERA5-WRFL showing a
 518 marked improvement (Fig. 9m).



519
 520 **Figure 9.** Annual mean near-surface atmospheric maximum temperature bias simulated over south-
 521 eastern Australia (WRF simulation inner domain) with respect to gridded observations for
 522 NARClIM2.0 RCMs forced by ERA-Interim for 2016 plus two months spin-up starting in November
 523 2015 (a-i), and corresponding NARClIM1.5 simulations for the same period forced by ERA5 (j-m).

524 In terms of RCM performances in simulating minimum temperature prior to switching the
 525 driving reanalyses, ERA-Interim-forced WRFJ-K-L RCMs of the CORDEX-CMIP5 era have lower
 526 overall biases for minimum temperature over the inner domain relative to the NARClIM2.0 ERA5-
 527 R1-R7 RCMs (i.e. ensemble mean absolute biases are 0.62K and 0.8K, respectively; Fig. 10b,j).
 528 However, the biases of each RCM generation vary geographically, such that the bias magnitudes for
 529 some ERA5-RCMs (e.g. R2-R3) are lower along coastal areas relative to ERA-Interim WRFJ-K-L
 530 over the same areas (Fig. 10d-e; k-m). Conversely, biases are lower over inland regions for ERA-
 531 Interim WRFJ-K-L relative to ERA5-RCMs.

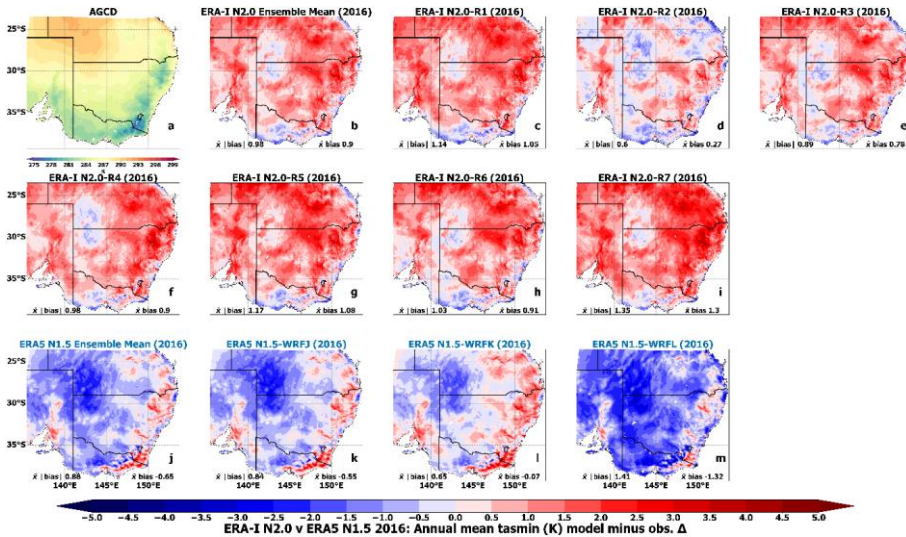
Deleted: |
 Deleted: |



534
 535 **Figure 10.** Annual mean near-surface atmospheric minimum temperature bias simulated over south-
 536 eastern Australia (WRF simulation inner domain) with respect to gridded observations for the period
 537 1981–2010 for NARcliM2.0 RCMs (b–i) and NARcliM1.5 RCMs (j–m). Stippling and panel
 538 boundary colouring as per Fig. 3.

539 Considering RCM simulations of mean minimum temperature with the driving reanalyses
 540 switched, performances are typically substantially poorer for the ERA5-forced WRFJ–K–L RCMs
 541 (Fig. 11) relative to their ERA-Interim-forced counterparts: the ensemble mean absolute biases are
 542 0.88K versus 0.62K, respectively. In contrast, although all NARcliM2.0 RCMs except R2 show
 543 performance degradations when forced with ERA-Interim instead of ERA5 (e.g. ensemble mean
 544 biases are 0.98K and 0.8K, respectively), these deteriorations are small (Fig. 11b–i).

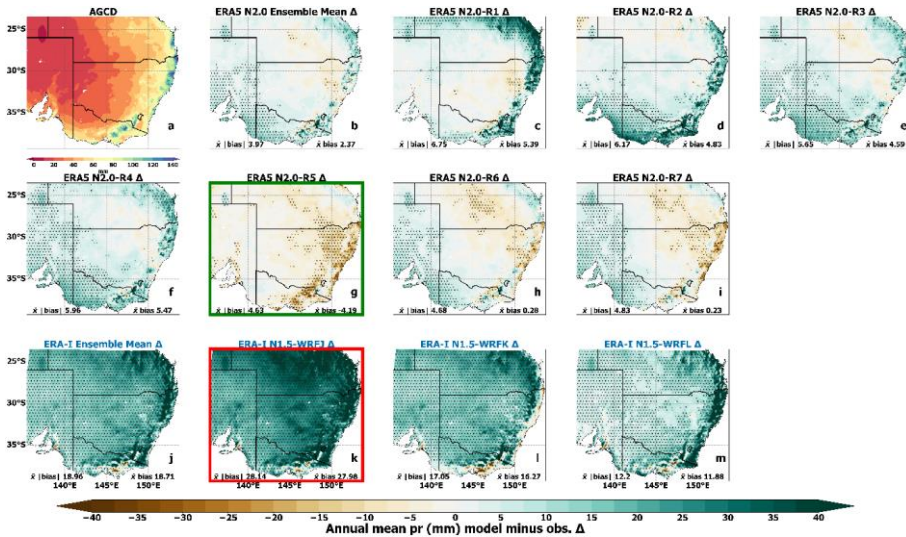
Deleted: |
 Deleted: |



547
 548 **Figure 11.** Annual mean near-surface atmospheric minimum temperature bias with respect to gridded
 549 observations for NARClIM2.0 RCMs forced by ERA-Interim for 2016 plus two months spin-up
 550 starting in November 2015 (a-i), and corresponding NARClIM1.5 simulations for the same period
 551 forced by ERA5 (j-m).

552 Improvements in the simulation of mean precipitation for ERA5-forced R1-R7 RCMs versus
 553 ERA-Interim WRFJ-K-L RCMs are especially evident over the high resolution south-eastern inner
 554 domain. At this scale, biases for several ERA5-forced R1-R7 RCMs are $< \sim 5$ mm compared to $> \sim 15$
 555 mm for the ERA-Interim-WRFJ-K-L RCMs (Fig. 12). Moreover, several improvements in the ERA5-
 556 RCM simulation of annual mean precipitation are apparent at convection permitting scale relative to
 557 over the 20 km outer domain. For instance, dry biases for ERA5-R3 and R5 along the eastern
 558 coastline are reduced at the convection-permitting scale.

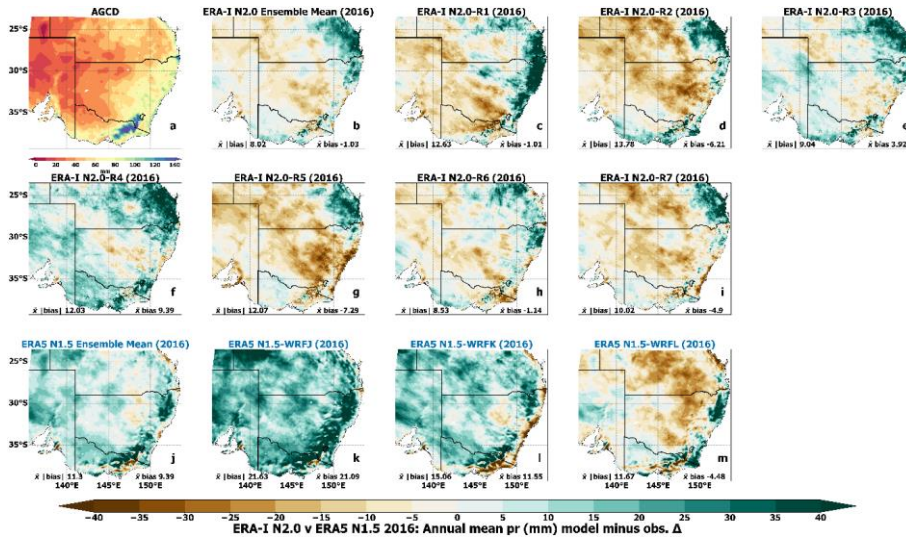
Deleted: the



560
 561 **Figure 12.** Annual mean precipitation bias simulated over south-eastern Australia (WRF simulation
 562 inner domain) with respect to gridded observations for the period 1981-2010 for NARClm2.0 RCMs
 563 (b-i) and NARClm1.5 RCMs (j-m). Stippling and panel boundary colouring as per Fig. 3.

564 Switching driving reanalyses and simulating annual mean precipitation produces results that
 565 show consistent, large changes in RCM performances when using the newer ERA5 data, versus ERA-
 566 Interim. Forcing the NARClm2.0 R1-R7 RCMs with ERA-Interim shows widespread marked
 567 increases in bias for annual mean precipitation for 2016 (Fig 13b-i) relative to the preceding
 568 simulations using ERA5, such that the ensemble area-averaged mean absolute bias deteriorates to 8.02
 569 mm versus 3.97 mm, i.e. roughly doubling the bias magnitude. Conversely, forcing WRFJ-K-L with
 570 ERA5 improves the simulation of annual mean precipitation with all RCMs showing reductions in
 571 bias (Fig. 13j-m), such that the ensemble mean absolute bias decreases from 18.96 mm to 11.3 mm.
 572 These performance improvements are smaller in magnitude as compared to the degradation in
 573 performance when switching the driving data for the NARClm2.0 R1-R7 RCMs.

- Deleted: and
- Deleted: bias
- Deleted: as compared
- Deleted: |
- Deleted: |
- Deleted: as compared to
- Deleted: small
- Deleted: |
- Deleted: |
- Deleted: 16.12



584
 585 **Figure 13.** Annual mean precipitation bias with respect to gridded observations for NARClIM2.0
 586 RCMs forced by ERA-Interim for 2016 plus two months spin-up starting in November 2015 (a-i), and
 587 corresponding NARClIM1.5 simulations for the same period forced by ERA5 (j-m).

588 **4. Discussion**

589 We have evaluated the capabilities of CORDEX-CMIP6 ERA5-driven RCMs in simulating the
 590 Australian climate and compared their performances to the previous generation of ERA-Interim
 591 forced RCMs produced for CORDEX-CMIP5. The newer generation of RCMs generally show
 592 improved simulations of maximum temperature and precipitation, but no improvements for minimum
 593 temperature. Several changes have been made to the design of the newer generation of RCMs,
 594 including different RCM physics parameterisations, model specifications, and the driving reanalysis is
 595 newer (ERA5). We found no evidence to suggest that the newer reanalysis contributes to the
 596 improvements in the simulation of maximum temperature by the ERA5 RCMs, whereas the opposite
 597 applies to the simulation of precipitation. This study focuses primarily on model evaluation with
 598 investigations of potential mechanisms underlying the varying performance profiles of the different
 599 RCM generations to be the subject of future research. This will be facilitated by the imminent
 600 publication of the NARClIM2.0 ERA5-RCM data.

601 **4.1 RCM performance evaluation**

602 As per the ERA-Interim driven RCMs, the NARClIM2.0 CORDEX-CMIP6 ERA5 RCMs are
 603 generally cold-biased for mean maximum temperature, however, their bias magnitudes are
 604 substantially lower relative to the CORDEX-CMIP5 ERA-Interim ensemble. The reductions in bias

605 magnitude for most CORDEX-CMIP6 ERA5-RCMs are especially marked for the convection-
606 permitting 4 km inner domain over south-eastern Australia. Similarly, these ERA5 RCMs show an
607 overall improved simulation of extreme maximum temperature over most of Australia relative to the
608 CORDEX-CMIP5 ERA-Interim forced RCMs. Improved simulation of mean and extreme maximum
609 temperature has important practical applications for climate impact assessment in Australia (e.g. Van
610 Oldenborgh et al., 2021; Di Virgilio et al., 2020a; Trancoso et al., 2020), as well as globally (e.g.
611 Vargas Zeppetello et al., 2022; Schleussner et al., 2016; Auffhammer et al., 2017).

612 Overall, CORDEX-CMIP6 ERA5-RCMs confer improvements in the simulation of mean
613 precipitation over [south-eastern](#) Australia relative to the CORDEX-CMIP5 ERA-Interim RCMs, with
614 two ERA5 RCMs in particular (R3, R4) showing considerable improvements [over this region](#).
615 Improvements in the simulation of mean precipitation by CORDEX-CMIP6 ERA5 RCMs are even
616 more marked at convection-permitting scale over south-eastern Australia, i.e. the ERA5 ensemble
617 mean is 3.97 mm versus 18.96 mm for the ERA-Interim ensemble. Given the significant impacts of
618 drought and floods in Australia (González Tánago et al., 2016; Gu et al., 2020), this improvement in
619 mean precipitation simulation is an encouraging result. The performance in simulating extreme
620 precipitation over the Australian continent is comparable between the CORDEX-CMIP6 ERA5
621 RCMs and most CORDEX-CMIP5 ERA-Interim RCMs, except WRFSSWA, CCAM and CCLM
622 which show strong biases. [However, at convection-permitting scale, some ERA5-RCMs show](#)
623 [improvements of around 10% in the simulation of extreme precipitation relative to the ERA-Interim](#)
624 [RCMs, except ERA5-R1 and R2 which are strongly wet-biased](#). For both mean and extreme
625 precipitation, ERA5 R1 and R2 are notable in that they are more wet-biased than the other ERA5
626 RCMs, especially over northern Australia [where all other ERA5-RCMs contain a systematic dry-bias](#).
627 The only physics parameterisation common to both ERA5-R1 and R2 is their use of WSM6
628 microphysics, and no other RCMs assessed here use this physics scheme, with ERA5-R3-R7 using
629 Thompson microphysics. A previous assessment of the performance of different WRF
630 parameterisations for a one-way nested inner domain over central Europe observed that WSM6
631 increases annual wet bias relative to other microphysical schemes tested, including the Thomson
632 scheme (Varga and Breuer, 2020). Notably, marked dry-biases over the monsoonal north for several
633 ERA5-forced RCMs correspond with warm maximum temperature biases over this region shown by
634 several ERA5 RCMs.

635 Whilst the ERA5 RCMs confer improvements to the simulation of maximum temperature and
636 precipitation relative to ERA-Interim models, the simulation of minimum temperature for all
637 timescales and statistics shows no improvement over the Australian continent. Focusing specifically
638 on the WRF RCM configurations in the ERA-Interim ensemble, WRFJ and WRFK simulate both
639 mean and extreme minimum temperature more accurately than the ERA5-forced models, though in
640 some cases the differences are minimal. The exception to the above result is that some ERA5-RCMs

Deleted: However, over the convection-permitting domain, many ERA5-RCMs show enhanced simulation of extreme precipitation relative to the ERA-Interim RCMs, except ERA5-R1 and R2 which are strongly wet-biased

645 simulate mean minimum temperature more accurately along south-eastern coastlines at the 4 km
646 convection-permitting scale.

647 **4.2 ERA5 versus ERA-Interim evaluations: potential implications for** 648 **CMIP6-forced dynamical downscaling**

649 It could be expected that differences in the reanalysis data sets used to force the two generations of
650 WRF RCM ensemble contribute to the varying RCM performance profiles observed. ERA5 is a more
651 recent reanalysis which comprises a range of improvements over ERA-Interim, for instance, increased
652 resolutions spanning horizontal (~31 km versus ~79 km), vertical (137 levels to 0.01 hPa versus 60 to
653 0.1 hPa), and temporal dimensions (hourly versus 6-hourly), among other features such as improved
654 parameterisations (Hersbach et al., 2020). ERA5 has been shown to confer improvements over ERA-
655 Interim in the simulation of processes such as convective updrafts, tropical cyclones, and other meso-
656 to synoptic-scale atmospheric features (Hoffmann et al., 2019) and in some cases the simulation of
657 rainfall (e.g. Nogueira, 2020). Our investigation into whether differences in the driving reanalyses
658 contribute to the varying RCM performances observed between the two WRF RCM ensembles
659 involved two assessments: i) comparisons of the ERA5 and ERA-Interim reanalyses against AGCD
660 observations to assess their degree of bias; ii) fourteen-month simulations where otherwise identically
661 parameterised NARClIM2.0 R1-R7 RCMs were forced by ERA-Interim as opposed to ERA5, and
662 similarly the WRFJ-K-L RCMs were forced with ERA5 instead of ERA-Interim.

663 Comparison of ERA5 and ERA-Interim reanalysis data versus observations for mean
664 maximum and minimum temperature and precipitation shows the expected results, i.e. that ERA5 data
665 are closer to observations relative to ERA-Interim for all variables, especially for mean precipitation.
666 Percentage differences in area-averaged mean absolute bias for annual means range from 25% for
667 minimum temperature to 65% for precipitation, also noting that performances during summer were
668 more divergent than at annual timescales. Therefore, in terms of the underlying reanalysis data used to
669 force the different WRF RCMs evaluated, ERA5 shows improvements relative to ERA-Interim.
670 Additionally, these improvements are of larger magnitude for mean precipitation than they are for
671 mean maximum and minimum temperature.

672 For the 1-year simulations where the driving reanalyses are switched, using ERA5 over ERA-
673 Interim gives a large performance improvement in the simulation of annual mean precipitation for the
674 CORDEX-CMIP5 WRFJ-K-L RCMs. In contrast, using ERA5 over ERA-Interim as the driving data
675 generally produces RCM performance degradations for both annual mean maximum and minimum
676 temperature. That is, a superior simulation of mean maximum and minimum temperature is generally
677 obtained for both generations of WRF RCM by using ERA-Interim instead of ERA5. These results
678 suggest that, at least for the different generations of WRF RCM assessed here in these 1-year
679 experiments, using a more accurate driving reanalysis for dynamical downscaling over this region

680 does not guarantee an enhanced simulation for all climatic variables. This result is surprising and
681 warrants further investigation. However, this finding suggests that the parameterisations and design
682 features of the WRF RCMs assessed play important roles in determining how well these RCMs
683 simulate mean maximum and minimum temperature. Consequently, the improved simulations of
684 maximum temperature by CORDEX-CMIP6 ERA5-RCMs relative to CORDEX-CMIP5 ERA-
685 Interim-RCMs are more attributable to model design choices, such as physics parameterisations
686 and/or improved resolution, rather than to the driving reanalyses per se. Additionally, that the
687 CORDEX-CMIP6 ERA5-forced R1-R7 RCMs do not improve the simulation of minimum
688 temperature relative to CORDEX-CMIP5 ERA-Interim-forced RCMs is not attributable to the change
689 from ERA-Interim to ERA5 as the driving reanalysis, rather, to aspect(s) of model
690 parameterisation/design. Conversely, substantial improvements in simulating mean precipitation by
691 CORDEX-CMIP6 ERA5-RCMs relative to CORDEX-CMIP5 ERA-Interim-forced RCMs appear (at
692 least in part) due to the improvements to the ERA5 driving reanalysis. There are limitations to these
693 comparative analyses switching the driving data, such as simulating for fourteen months and not a
694 climatological period. Nevertheless, the present evaluations suggest that whether CORDEX-CMIP6
695 dynamical downscaling of CMIP6 GCMs produces improved regional climate simulations relative to
696 CORDEX-CMIP5 downscaling may depend in large part, at least for some variables/statistics, on
697 RCM parameterisations and other design choices. However, the generality of these findings to other
698 RCM types, configurations, study domains, and downscaling experiments warrants further research as
699 these results may be specific to the WRF RCMs and domains assessed here.

700 **4.3 ERA5-R1-R7 and CMIP6-forced dynamical downscaling**

701 Although a single 'all-round' best-performing ERA5-RCM configuration cannot be selected, the RCM
702 performances for the climate variables and statistics assessed here yield some insights if selecting a
703 subset of ERA5-RCM configurations for subsequent CMIP6-forced downscaling. Overall, ERA5-R1
704 provides a good simulation of both mean and extreme maximum temperature and is broadly
705 comparable to the other ERA5-RCMs with respect to minimum temperature. However, its simulation
706 of mean and extreme precipitation is relatively poor as compared to most ERA5-RCMs. ERA5-R2 has
707 an unusual performance profile relative to the other ERA5-RCMs. Although ERA5-R2 shows
708 generally good performance for minimum temperature, extreme maximum temperature and
709 precipitation, it shows poor performance for mean maximum temperature in that is considerably more
710 cold-biased than the other ERA5-RCMs. ERA5-R2 is the only ERA5-forced RCM configuration in
711 this ensemble to use Kain-Fritsch cumulus physics, and it shows mean maximum temperature biases
712 of roughly similar magnitude and spatial pattern as the ERA-Interim WRFJ and WRFK RCMs which
713 also use the same scheme. However, ERA5-R2 also generates a strong mean maximum temperature
714 cold bias over south-eastern Australia at the 4 km convection-permitting scale which does not use

715 cumulus parameterisation. ERA5-R3 shows good performance for mean minimum temperature and
716 mean precipitation and reasonable performance for mean maximum temperature. The performance of
717 ERA5-R4 is broadly similar to ERA5-R3, but it has substantially inferior performance versus ERA5-
718 R3 for maximum and minimum temperature extremes. ERA5-R5 shows consistently good
719 performance for maximum temperature. The performance of ERA5-R5 in simulating precipitation
720 over Australia at 20 km resolution is not impressive versus the other ERA5-RCMs and it shows strong
721 dry biases over northern Australia. However, ERA5-R5 is the best-performing model in this ensemble
722 for mean precipitation at the 4 km convection permitting scale over south-eastern Australia. Both
723 ERA5-R6 and ERA5-R7 frequently show the strongest biases, typically over large regions such as
724 eastern Australia for both temperature variables, and over northern Australia for precipitation. As
725 such, they are the poorest performers overall in this ERA5 ensemble, with performance for extreme
726 minimum temperature often being particularly poor.

727 From the specific perspective of the ERA5-RCM performances, and based on the present
728 evaluations, overall ERA5-R3 and ERA5-R5 may be considered favourable RCM configurations for
729 CMIP6-forced dynamical downscaling. However, as noted, some other ERA5 RCM configurations
730 show good performance for specific variables and statistics, and thus could warrant inclusion in a
731 larger ensemble and/or one adopting a sparse matrix approach (Christensen and Kjellström, 2020).

732 **5. Conclusions**

733 This study forms the first part of a series of simulations for the CORDEX Australasia domain,
734 wherein we document model performances of ERA5 reanalysis-forced RCMs, and this is the first set
735 of simulations as required by the CORDEX-CMIP6 framework. We compared our results against
736 ERA-Interim driven simulations which was part of the CORDEX-CMIP5 framework. While model
737 versions and physics options were different between these two generations of reanalysis-forced RCM
738 simulations, overall, our results show the NARClIM2.0 ERA5-forced RCMs confer improved
739 simulations for maximum temperature and precipitation, but not for minimum temperature.

740 The simulation of precipitation by the NARClIM2.0 RCMs show several improvements at the
741 4 km convection permitting scale relative to the 20 km outer domain. For example, dry biases are
742 reduced for the convection-permitting domain where convection is represented explicitly, relative to
743 the 20 km outer domain which uses a convective parametrisation. Convection schemes can be a
744 source of deficiencies in RCM simulations of precipitation (e.g. Jones and Randall, 2011). It may be
745 expected that the improved representation of convection for the 4 km domain may positively
746 influence the simulation of high-impact phenomena such as short-duration precipitation extremes.
747 Nevertheless, our results for the CORDEX-Australasia domain suggest that the choice of
748 microphysics scheme is important, especially for precipitation extremes.

Deleted: at such scales

750 Whilst ERA5 reanalysis data show better representations of the observed Australian climate
751 than ERA-Interim, only improvements in the simulation of mean precipitation by the CORDEX-
752 CMIP6 ERA5-RCMs appear at least partly attributable to the increased accuracy of ERA5 driving
753 reanalyses. Conversely, the change in driving reanalysis from ERA-Interim to ERA5 is not a major
754 factor underlying improvements in the simulation of maximum temperature by the CORDEX-CMIP6
755 RCMs assessed, suggesting that their performance improvements are more attributable to changes in
756 RCM parameterisation and design. The different land surface schemes (e.g. Noah-Unified versus
757 Noah-MP) likely play a role in RCM skill in simulating maximum temperature, as well as changing
758 the land surface feedback (via soil moisture) to the simulation of precipitation – these possibilities
759 require more extensive analysis to investigate. Equally, differences in the underlying driving reanalyses
760 do not explain the absence of overall improvements in the simulation of minimum temperature by the
761 newer CORDEX-CMIP6 RCMs. It is important to be cautious of generalising the present results to
762 other regions globally, as region-specific RCM optimisation is necessary.

763 Our present focus was to evaluate the performances of the different RCM generations
764 assessed here. Future work will explore other topics, such as the potential influences of the different
765 RCM physics configurations and their associated biases on the nature of the future change signals in
766 subsequent CMIP6 GCM-forced simulations, e.g. when holding the driving GCM data constant.
767 Additionally, future model-intercomparison studies that compare biases between the different RCMs
768 contributing to CORDEX-Australasia will be valuable.

769 Results presented here are relevant for other CORDEX-CMIP6/CORDEX2 modelling
770 projects. Maximum temperature and precipitation are important inputs to climate impact assessments
771 in Australia, and globally. The improvements in simulating maximum temperature and precipitation
772 conferred by CORDEX-CMIP6 ERA5-forced RCMs evaluated here indicate that using a subset of the
773 RCMs in this ensemble for future CMIP6-forced downscaling over CORDEX Australasia could yield
774 benefits in simulating regional climate.

775 6. Code Availability

776 The Weather Research and Forecasting (WRF) version 4.1.2 and all model configuration files used
777 in this study are available on Zenodo at: <https://doi.org/10.5281/zenodo.11189898>

778 7. Data Availability

779 Data for the seven CORDEX-CMIP6 ERA5-forced R1-R7 RCMs are being made available via
780 [National Computing Infrastructure](#) (NCI). WRF namelist settings for the CORDEX-CMIP6 ERA5-
781 forced RCMs R1-R7 are shown in Supplementary Material [Fig. S32](#). Data for the three ERA-Interim
782 forced WRFJ-K-L RCMs are available via the [New South Wales Climate Data Portal](#) and [CORDEX-](#)

Deleted: the

Deleted: o

Deleted: of

Deleted: 20

787 [DKRZ](#), and data for ERA-Interim forced CCAM, CCLM and WRFSSWA are available via
788 [CORDEX-DKRZ](#).

789 **8. Author Contribution**

790 GDV and JPE designed the models and the simulations. FJ, ET, and CT setup the models and
791 conducted the model simulations with contributions from JPE, JK, JA, and YL. GDV prepared the
792 manuscript with contributions from all co-authors.

793 **9. Competing Interests**

794 The authors declare that they have no conflict of interest, noting that JK is a Topic Editor of
795 Geoscientific Model Development.

796 **10. Funding**

797 This research was supported by the New South Wales Department of Planning and Environment as
798 part of the NARClIM2.0 dynamical downscaling project contributing to CORDEX Australasia.
799 Funding was provided by the NSW Climate Change Fund for NSW and ACT Regional Climate
800 Modelling (NARClIM) Project. This research was undertaken with the assistance of resources and
801 services from the National Computational Infrastructure (NCI), which is supported by the Australian
802 Government.

803 Jason P. Evans acknowledges the support of the Australian Research Council Centre of Excellence for
804 Climate Extremes (CE170100023) and the Climate Systems Hub of the Australian Governments
805 National Environmental Science Program.

806 **11. References**

- 807 Auffhammer, M., Baylis, P., and Hausman, C. H.: Climate change is projected to have severe impacts
808 on the frequency and intensity of peak electricity demand across the United States,
809 Proceedings of the National Academy of Sciences, 114, 1886-1891,
810 doi:10.1073/pnas.1613193114, 2017.
- 811 Bucchignani, E., Mercogliano, P., Rianna, G., and Panitz, H. J.: Analysis of ERA-Interim-driven
812 COSMO-CLM simulations over Middle East - North Africa domain at different spatial
813 resolutions, International Journal of Climatology, 36, 3346-3369, 10.1002/joc.4559, 2016.
- 814 Bureau of Meteorology: Australian Gridded Climate Data (AGCD) ; v2.0.0 Snapshot (1900-01-01 to
815 2020-05-31), 2020.
- 816 Bureau of Meteorology, Annual climate statement 2016, last access: February 15 2024.

817 Chou, M. D., Suarez, M. J., Liang, X. Z., and Yan, M. M. H.: A thermal infrared radiation
818 parameterization for atmospheric studies, NASA Tech. Memo. NASA/TM-2001-104606, 19,
819 68 pp. <https://ntrs.nasa.gov/citations/20010072848>, 2001.

820 Christensen, O. B. and Kjellström, E.: Partitioning uncertainty components of mean climate and
821 climate change in a large ensemble of European regional climate model projections, *Climate*
822 *Dynamics*, 54, 4293-4308, 10.1007/s00382-020-05229-y, 2020.

823 Dee, D. P., Uppala, S. M., Simmons, A. J., Berrisford, P., Poli, P., Kobayashi, S., Andrae, U.,
824 Balmaseda, M. A., Balsamo, G., Bauer, P., Bechtold, P., Beljaars, A. C. M., van de Berg, L.,
825 Bidlot, J., Bormann, N., Delsol, C., Dragani, R., Fuentes, M., Geer, A. J., Haimberger, L.,
826 Healy, S. B., Hersbach, H., Hólm, E. V., Isaksen, I., Kållberg, P., Köhler, M., Matricardi, M.,
827 McNally, A. P., Monge-Sanz, B. M., Morcrette, J. J., Park, B. K., Peubey, C., de Rosnay, P.,
828 Tavolato, C., Thépaut, J. N., and Vitart, F.: The ERA-Interim reanalysis: configuration and
829 performance of the data assimilation system, *Quarterly Journal of the Royal Meteorological*
830 *Society*, 137, 553-597, 10.1002/qj.828, 2011.

831 Di Luca, A., de Elia, R., and Laprise, R.: Potential for added value in precipitation simulated by high-
832 resolution nested Regional Climate Models and observations, *Clim. Dyn.*, 38, 1229-1247,
833 10.1007/s00382-011-1068-3, 2012.

834 Di Virgilio, G., Evans, J. P., Clarke, H., Sharples, J., Hirsch, A. L., and Hart, M. A.: Climate Change
835 Significantly Alters Future Wildfire Mitigation Opportunities in Southeastern Australia,
836 *Geophys. Res. Lett.*, 47, e2020GL088893, <https://doi.org/10.1029/2020GL088893>, 2020a.

837 Di Virgilio, G., Evans, J. P., Di Luca, A., Grose, M. R., Round, V., and Thatcher, M.: Realised added
838 value in dynamical downscaling of Australian climate change, *Clim. Dyn.*, 54, 4675-4692,
839 10.1007/s00382-020-05250-1, 2020b.

840 Di Virgilio, G., Evans, J. P., Di Luca, A., Olson, R., Argüeso, D., Kala, J., Andrys, J., Hoffmann, P.,
841 Katzfey, J. J., and Rockel, B.: Evaluating reanalysis-driven CORDEX regional climate
842 models over Australia: model performance and errors, *Clim. Dyn.*, 53, 2985-3005,
843 10.1007/s00382-019-04672-w, 2019.

844 Di Virgilio, G., Evans, J., Ji, F., Tam, E., Kala, J., Andrys, J., Thomas, C., Choudhury, D., Rocha, C.,
845 White, S., Li, Y., El Rafei, M., Goyal, R., Riley, M., and Lingala, J.: Design, evaluation and
846 future projections of the NARClIM2.0 CORDEX-CMIP6 Australasia regional climate
847 ensemble, *Geosci. Model Dev. Discuss.*, 2024, 1-48, <https://doi.org/10.5194/gmd-2024-87>,
848 (in review).

849 Ekström, M., Grose, M. R., and Whetton, P. H.: An appraisal of downscaling methods used in climate
850 change research, *Wiley Interdisciplinary Reviews: Climate Change*, 6, 301-319,
851 10.1002/wcc.339, 2015.

852 Evans, A., Jones, D., Lelley, S., and Smalley, R.: An Enhanced Gridded Rainfall Analysis Scheme
853 for Australia, Australian Bureau of Meteorology, 2020.

854 Evans, J. P., Ji, F., Lee, C., Smith, P., Argüeso, D., and Fita, L.: Design of a regional climate
855 modelling projection ensemble experiment - NARCLiM, *Geosci. Model Dev.*, 7, 621-629,
856 10.5194/gmd-7-621-2014, 2014.

857 Eyring, V., Bony, S., Meehl, G. A., Senior, C. A., Stevens, B., Stouffer, R. J., and Taylor, K. E.:
858 Overview of the Coupled Model Intercomparison Project Phase 6 (CMIP6) experimental
859 design and organization, *Geosci. Model Dev.*, 9, 1937-1958, 10.5194/gmd-9-1937-2016,
860 2016.

861 Giorgi, F.: Regional climate modeling: Status and perspectives, *J. Phys. IV*, 139, 101-118,
862 10.1051/jp4:2006139008, 2006.

863 Giorgi, F.: Thirty Years of Regional Climate Modeling: Where Are We and Where Are We Going
864 next?, *Journal of Geophysical Research: Atmospheres*, 124, 5696-5723,
865 10.1029/2018jd030094, 2019.

866 Giorgi, F. and Bates, G. T.: The Climatological Skill of a Regional Model over Complex Terrain,
867 *Monthly Weather Review*, 117, 2325-2347, 10.1175/1520-0493(1989)117, 1989.

868 Giorgi, F., Jones, C., and Asrar, G.: Addressing climate information needs at the regional level: The
869 CORDEX framework, *WMO Bulletin*, 53, 175-183, 2009.

870 González Tánago, I., Urquijo, J., Blauhut, V., Villarroya, F., and De Stefano, L.: Learning from
871 experience: a systematic review of assessments of vulnerability to drought, *Nat. Hazards*, 80,
872 951-973, 10.1007/s11069-015-2006-1, 2016.

873 Grose, M. R., Narsey, S., Delage, F. P., Dowdy, A. J., Bador, M., Boschhat, G., Chung, C., Kajtar, J.
874 B., Rauniyar, S., Freund, M. B., Lyu, K., Rashid, H., Zhang, X., Wales, S., Trenham, C.,
875 Holbrook, N. J., Cowan, T., Alexander, L., Arblaster, J. M., and Power, S.: Insights From
876 CMIP6 for Australia's Future Climate, *Earth's Future*, 8, e2019EF001469,
877 <https://doi.org/10.1029/2019EF001469>, 2020.

878 Grose, M., Narsey, S., Tranco, R., Mackallah, C., Delage, F., Dowdy, A., Di Virgilio, G.,
879 Watterson, I., Dobrohotoff, P., Rashid, H. A., Rauniyar, S., Henley, B., Thatcher, M., Syktus,
880 J., Abramowitz, G., Evans, J. P., Su, C.-H., and Takbash, A.: A CMIP6-based multi-model
881 downscaling ensemble to underpin climate change services in Australia, *Climate Services*, 30,
882 100368, <https://doi.org/10.1016/j.cliser.2023.100368>, 2023.

883 Gu, X. H., Zhang, Q., Li, J. F., Liu, J. Y., Xu, C. Y., and Sun, P.: The changing nature and projection
884 of floods across Australia, *J. Hydrol.*, 584, 10.1016/j.jhydrol.2020.124703, 2020.

885 Gutowski, W. J., Giorgi, F., Timbal, B., Frigon, A., Jacob, D., Kang, H. S., Raghavan, K., Lee, B.,
886 Lennard, C., Nikulin, G., O'Rourke, E., Rixen, M., Solman, S., Stephenson, T., and Tangang,
887 F.: WCRP COordinated Regional Downscaling EXperiment (CORDEX): a diagnostic MIP
888 for CMIP6, *Geoscientific Model Development*, 9, 4087-4095, 10.5194/gmd-9-4087-2016,
889 2016.

890 Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J.,
891 Peubey, C., Radu, R., Schepers, D., Simmons, A., Soci, C., Abdalla, S., Abellan, X.,
892 Balsamo, G., Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., De Chiara, G., Dahlgren, P.,
893 Dee, D., Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer, A.,
894 Haimberger, L., Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux, P.,
895 Lopez, P., Lupu, C., Radnoti, G., de Rosnay, P., Rozum, I., Vamborg, F., Villaume, S., and
896 Thépaut, J.-N.: The ERA5 global reanalysis, *Quarterly Journal of the Royal Meteorological*
897 *Society*, 146, 1999-2049, <https://doi.org/10.1002/qj.3803>, 2020.

898 Hoffmann, L., Gunther, G., Li, D., Stein, O., Wu, X., Griessbach, S., Heng, Y., Konopka, P., Muller,
899 R., Vogel, B., and Wright, J. S.: From ERA-Interim to ERA5: the considerable impact of
900 ECMWF's next-generation reanalysis on Lagrangian transport simulations, *Atmos. Chem.*
901 *Phys.*, 19, 3097-3124, 10.5194/acp-19-3097-2019, 2019.

902 Hong, S.-Y., Noh, Y., and Dudhia, J.: A New Vertical Diffusion Package with an Explicit Treatment
903 of Entrainment Processes, *Monthly Weather Review*, 134, 2318-2341,
904 <https://doi.org/10.1175/MWR3199.1>, 2006.

905 Hong, S. Y. and Lim, J.-O. J.: The WRF Single-Moment 6-Class Microphysics Scheme (WSM6),
906 *Asia-Pac. J. Atmos. Sci.*, 42, 129-151, 2006.

907 Howard, E., Su, C. H., Stassen, C., Naha, R., Ye, H., Pepler, A., Bell, S. S., Dowdy, A. J., Tucker, S.
908 O., and Franklin, C.: Performance and process-based evaluation of the BARPA-R
909 Australasian regional climate model version 1, *Geosci. Model Dev.*, 17, 731-757,
910 10.5194/gmd-17-731-2024, 2024.

911 Hsiang, S., Kopp, R., Jina, A., Rising, J., Delgado, M., Mohan, S., Rasmussen, D. J., Muir-Wood, R.,
912 Wilson, P., Oppenheimer, M., Larsen, K., and Houser, T.: Estimating economic damage from
913 climate change in the United States, *Science*, 356, 1362-1368, 10.1126/science.aal4369, 2017.

914 Iacono, M. J., Delamere, J. S., Mlawer, E. J., Shephard, M. W., Clough, S. A., and Collins, W. D.:
915 Radiative forcing by long-lived greenhouse gases: Calculations with the AER radiative
916 transfer models, *Journal of Geophysical Research: Atmospheres*, 113,
917 <https://doi.org/10.1029/2008JD009944>, 2008.

918 IPCC: Climate Change 2021: The Physical Science Basis. Contribution of Working Group I to the
919 Sixth Assessment Report of the Intergovernmental Panel on Climate Change, Cambridge
920 University Press, 2021.

921 Janjić, Z. I.: Comments on "Development and Evaluation of a Convection Scheme for Use in Climate
922 Models", *Journal of the Atmospheric Sciences*, 57, 3686-3686, [https://doi.org/10.1175/1520-0469\(2000\)057<3686:CODAEO>2.0.CO;2](https://doi.org/10.1175/1520-0469(2000)057<3686:CODAEO>2.0.CO;2), 2000.

924 Jones, T. R. and Randall, D. A.: Quantifying the limits of convective parameterizations, *Journal of*
925 *Geophysical Research: Atmospheres*, 116, <https://doi.org/10.1029/2010JD014913>, 2011.

926 Kain, J. S.: The Kain-Fritsch convective parameterization: An update, *Journal of Applied*
927 *Meteorology*, 43, 170-181, 10.1175/1520-0450, 2004.

928 Laprise, R.: Regional climate modelling, *J. Comput. Phys.*, 227, 3641-3666,
929 10.1016/j.jcp.2006.10.024, 2008.

930 Ma, M. N., Ou, T. H., Liu, D. Q., Wang, S. Y., Fang, J., and Tang, J. P.: Summer regional climate
931 simulations over Tibetan Plateau: from gray zone to convection permitting scale, *Clim. Dyn.*,
932 10.1007/s00382-022-06314-0, 2022.

933 McGregor, J. L. and Dix, M. R.: An updated description of the Conformal-Cubic atmospheric model,
934 *High Resolution Numerical Modelling of the Atmosphere and Ocean*, Springer, New York,
935 51-75 pp., 10.1007/978-0-387-49791-4_4, 2008.

936 Nakanishi, M. and Niino, H.: Development of an Improved Turbulence Closure Model for the
937 Atmospheric Boundary Layer, *Journal of the Meteorological Society of Japan. Ser. II*, 87,
938 895-912, 10.2151/jmsj.87.895, 2009.

939 Nishant, N., Evans, J. P., Di Virgilio, G., Downes, S. M., Ji, F., Cheung, K. K. W., Tam, E., Miller, J.,
940 Beyer, K., and Riley, M. L.: Introducing NARCLiM1.5: Evaluating the Performance of
941 Regional Climate Projections for Southeast Australia for 1950–2100, *Earth's Future*, 9,
942 e2020EF001833, <https://doi.org/10.1029/2020EF001833>, 2021.

943 Niu, G.-Y., Yang, Z.-L., Mitchell, K. E., Chen, F., Ek, M. B., Barlage, M., Kumar, A., Manning, K.,
944 Niyogi, D., Rosero, E., Tewari, M., and Xia, Y.: The community Noah land surface model
945 with multiparameterization options (Noah-MP): 1. Model description and evaluation with
946 local-scale measurements, *Journal of Geophysical Research: Atmospheres*, 116,
947 10.1029/2010jd015139, 2011.

948 Nogueira, M.: Inter-comparison of ERA-5, ERA-interim and GPCP rainfall over the last 40 years:
949 Process-based analysis of systematic and random differences, *J. Hydrol.*, 583,
950 10.1016/j.jhydrol.2020.124632, 2020.

951 Panitz, H.-J., Dosio, A., Büchner, M., Lüthi, D., and Keuler, K.: COSMO-CLM (CCLM) climate
952 simulations over CORDEX-Africa domain: analysis of the ERA-Interim driven simulations at
953 0.44° and 0.22° resolution, *Clim. Dyn.*, 42, 3015-3038, 10.1007/s00382-013-1834-5, 2014.

954 Perkins, S. E., Pitman, A. J., Holbrook, N. J., and McAneney, J.: Evaluation of the AR4 climate
955 models' simulated daily maximum temperature, minimum temperature, and precipitation over
956 Australia using probability density functions, *J. Clim.*, 20, 4356-4376, 10.1175/jcli4253.1,
957 2007.

958 Pleim, J. E.: A Combined Local and Nonlocal Closure Model for the Atmospheric Boundary Layer.
959 Part I: Model Description and Testing, *J. Appl. Meteorol. Climatol.*, 46, 1383-1395,
960 <https://doi.org/10.1175/JAM2539.1>, 2007.

961 Reder, A., Raffa, M., Padulano, R., Rianna, G., and Mercogliano, P.: Characterizing extreme values
962 of precipitation at very high resolution: An experiment over twenty European cities, *Weather*
963 and *Climate Extremes*, 35, 10.1016/j.wace.2022.100407, 2022.

964 Rockel, B., Will, A., and Hense, A.: The Regional Climate Model COSMO-CLM(CCLM), *Meteorol.*
965 *Z.*, 17, 347-348, 10.1127/0941-2948/2008/0309, 2008.

966 Schleussner, C. F., Lissner, T. K., Fischer, E. M., Wohland, J., Perrette, M., Golly, A., Rogelj, J.,
967 Childers, K., Schewe, J., Frieler, K., Mengel, M., Hare, W., and Schaeffer, M.: Differential
968 climate impacts for policy-relevant limits to global warming: the case of 1.5 °C and 2 °C,
969 *Earth Syst. Dynam.*, 7, 327-351, 10.5194/esd-7-327-2016, 2016.

970 Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Barker, D. M., Wang, W., and Powers, J. G.:
971 A description of the Advanced Research WRF Version 3. NCAR Tech Note NCAR/TN-
972 475+STR. NCAR, Boulder, CO, 2008.

973 Stouffer, R. J., Eyring, V., Meehl, G. A., Bony, S., Senior, C., Stevens, B., and Taylor, K. E.: CMIP5
974 Scientific Gaps and Recommendations for CMIP6, *Bulletin of the American Meteorological*
975 *Society*, 98, 95-105, 10.1175/bams-d-15-00013.1, 2017.

976 Tebaldi, C., Arblaster, J. M., and Knutti, R.: Mapping model agreement on future climate projections,
977 *Geophys. Res. Lett.*, 38, doi:10.1029/2011GL049863, 2011.

978 Tewari, M., Wang, W., Dudhia, J., LeMone, M. A., Mitchell, K., Ek, M., Gayno, G., Wegiel, J., and
979 Cuenca, R.: Implementation and verification of the united NOAA land surface model in the
980 WRF model, 11-15 pp.2016.

981 Thompson, G., Field, P. R., Rasmussen, R. M., and Hall, W. D.: Explicit Forecasts of Winter
982 Precipitation Using an Improved Bulk Microphysics Scheme. Part II: Implementation of a
983 New Snow Parameterization, *Monthly Weather Review*, 136, 5095-5115,
984 <https://doi.org/10.1175/2008MWR2387.1>, 2008.

985 Tiedtke, M.: A Comprehensive Mass Flux Scheme for Cumulus Parameterization in Large-Scale
986 Models, *Monthly Weather Review*, 117, 1779-1800, 10.1175/1520-
987 0493(1989)117<1779:acmfsf>2.0.co;2, 1989.

988 Torma, C., Giorgi, F., and Coppola, E.: Added value of regional climate modeling over areas
989 characterized by complex terrain—Precipitation over the Alps, *Journal of Geophysical*
990 *Research: Atmospheres*, 120, 3957-3972, 10.1002/2014JD022781, 2015.

991 Trancoso, R., Syktus, J., Toombs, N., Ahrens, D., Wong, K. K.-H., and Pozza, R. D.: Heatwaves
992 intensification in Australia: A consistent trajectory across past, present and future, *Sci. Total*
993 *Environ.*, 742, 140521, <https://doi.org/10.1016/j.scitotenv.2020.140521>, 2020.

994 Van de Walle, J., Thiery, W., Brousse, O., Souverijns, N., Demuzere, M., and van Lipzig, N. P. M.: A
995 convection-permitting model for the Lake Victoria Basin: evaluation and insight into the
996 mesoscale versus synoptic atmospheric dynamics, *Clim. Dyn.*, 54, 1779-1799,
997 10.1007/s00382-019-05088-2, 2020.

998 van Oldenborgh, G. J., Krikken, F., Lewis, S., Leach, N. J., Lehner, F., Saunders, K. R., van Weele,
999 M., Haustein, K., Li, S., Wallom, D., Sparrow, S., Arrighi, J., Singh, R. K., van Aalst, M. K.,
1000 Philip, S. Y., Vautard, R., and Otto, F. E. L.: Attribution of the Australian bushfire risk to
1001 anthropogenic climate change, *Nat. Hazards Earth Syst. Sci.*, 21, 941-960, 10.5194/nhess-21-
1002 941-2021, 2021.

1003 Varga, A. J. and Breuer, H.: Sensitivity of simulated temperature, precipitation, and global radiation
1004 to different WRF configurations over the Carpathian Basin for regional climate applications,
1005 *Clim. Dyn.*, 55, 2849-2866, 10.1007/s00382-020-05416-x, 2020.

1006 Vargas Zeppetello, L. R., Raftery, A. E., and Battisti, D. S.: Probabilistic projections of increased heat
1007 stress driven by climate change, *Communications Earth & Environment*, 3, 183,
1008 10.1038/s43247-022-00524-4, 2022.

1009 Zhou, X., Yang, K., Ouyang, L., Wang, Y., Jiang, Y. Z., Li, X., Chen, D. L., and Prein, A.: Added
1010 value of kilometer-scale modeling over the third pole region: a CORDEX-CPTP pilot study,
1011 *Clim. Dyn.*, 57, 1673-1687, 10.1007/s00382-021-05653-8, 2021.