

# Lambda-PFLOTRAN 1.0: Workflow for Incorporating Organic Matter Chemistry Informed by Ultra High Resolution Mass Spectrometry into Biogeochemical Modeling

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**Abstract.** Organic matter (OM) composition plays a central role in microbial respiration of dissolved organic matter and subsequent biogeochemical reactions. Here, a direct connection of organic matter chemistry and thermodynamics to reactive transport simulators has been achieved through the newly developed Lambda-PFLOTRAN workflow tool that succinctly incorporates carbon chemistry data generated from Fourier transform ion cyclotron resonance mass spectrometry (FTICR-MS) into reaction networks to simulate organic matter degradation and the resulting biogeochemistry. Lambda-PFLOTRAN is a python-based workflow, executed through a Jupyter Notebook interface, that digests raw FTICR-MS data, develops a representative reaction network based on substrate-explicit thermodynamic modeling (also termed lambda modeling due to its key thermodynamic parameter  $\lambda$  used therein), and completes a biogeochemical simulation with the open source, reactive flow and transport code PFLOTRAN. The workflow consists of the following five steps: configuration, thermodynamic (lambda) analysis, sensitivity analysis, parameter estimation, and simulation output and visualization. Two test cases are provided to demonstrate the functionality of the Lambda-PFLOTRAN workflow. The first test case uses laboratory incubation data of temporal oxygen depletion to fit lambda parameters (i.e., maximum utilization rate and microbial carrying capacity). A slightly more complex second test case fits multiple lambda formulation and soil organic matter release parameters to temporal greenhouse gas generation measured during a soil incubation. Overall, the Lambda-PFLOTRAN workflow facilitates upscaling by using molecular-scale characterization to inform biogeochemical processes occurring at larger scales.

## 30 **1 Introduction**

31 Microbial respiration of dissolved organic carbon (DOC) is a main driver of environmental biogeochemical processes.  
32 Mechanistic biogeochemical models often rely on lumping organic matter into a few distinct carbon pools (e.g.,  
33 dissolved, sorbed, mineral associated or refractory, labile, etc.) (e.g., Fatichi, et al., 2019, Robertson et al., 2019, Wang  
34 et al., 2013) but do not fully consider the properties of the organic matter (OM) compounds individually. Pooled  
35 carbon approaches have benefits, such as assigning variable levels of bioavailability, however, this approach does not  
36 capture the complex temporal dynamics of respiration driven by OM composition, as aerobic respiration rates have  
37 been linked to organic carbon concentration, thermodynamics of the OM (Stegen et al., 2018, Garayburu-Caruso et  
38 al., 2020), as well as the diversity of OM compounds present (Lehmann et al. 2020, Stegen et al., 2022). Such findings  
39 highlight the importance of incorporating individual OM chemistry into biogeochemical modeling to capture, and  
40 ultimately predict, system behavior more accurately.

41 There are many advanced instrumentation techniques capable of detecting and identifying individual OM formulae  
42 that comprise a bulk OM sample (e.g., GC-MS, HPLC-MS, Fourier transform ion cyclotron resonance mass  
43 spectrometry [FTICR-MS], etc.). For instance, FTICR-MS is a powerful, high-resolution, method that identifies  
44 molecular formulae for individual organic compounds. In any given environmental sample, FTICR-MS (or other ultra  
45 high-resolution methods) will typically resolve thousands of discrete OM molecular formulae, each with a unique  
46 mass and elemental composition (Cooper et al., 2020, Bahureksa et al., 2021). However, untargeted analytical  
47 techniques like FTICR-MS are only able to determine if a compound is present and cannot quantify the total  
48 concentration associated with each organic matter molecule. Still, such techniques do provide immense amounts of  
49 characterization data encompassing a deeper analytical window than measuring a small number of individual  
50 biomarkers quantitatively (e.g., Ward et al., 2013). Utilizing such high-resolution molecular data in reactive transport  
51 modeling frameworks affords new opportunity to advance carbon cycling in terrestrial, riverine and coastal systems  
52 despite of various theoretical and computational challenges.

53 Substrate-explicit thermodynamic modeling (SXTM) provides an avenue for incorporating individual OM reactivity  
54 based on thermodynamics (Song et al., 2020) into reactive transport models. The SXTM procedure takes the individual  
55 chemical formula derived from FTICR-MS (or another high-resolution technique) and uses its thermodynamic  
56 properties to generate an oxidation reaction for each molecular formula present in a sample. The corresponding  
57 reaction stoichiometry is then determined by considering catabolic, anabolic, and metabolic reactions and balancing  
58 energy for the overall metabolic reaction, allowing for the development of an aerobic respiration expression for each  
59 OM formula.

60 Still, the sheer number of compounds identified in each sample proves difficult for model integration. Typically,  
61 reactive transport simulators consider only a small number of primary species in their reaction networks, and most  
62 could not support modeling each of the thousands of organic matter molecules individually. Here, the developed  
63 Lambda-PFLOTRAN workflow addresses this challenge through grouping, or binning, similar compounds based on

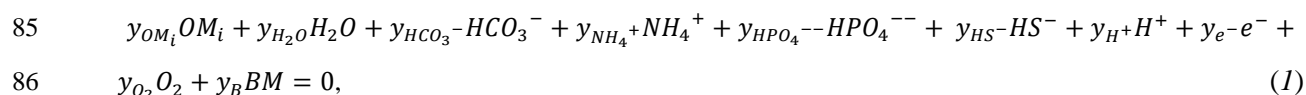
64 their thermodynamic properties, allowing for the number of species considered within the reaction network to be  
65 reduced, and thus decreasing the required computational resources.

66 Lambda-PFLOTRAN is a python-based workflow that digests raw FTICR-MS data, develops a representative reaction  
67 network based on substrate-explicit thermodynamic modeling (Song et al., 2020), and completes a biogeochemical  
68 simulation with the open source, parallel reactive flow and transport code, PFLOTRAN (Hammond et al., 2014).  
69 PFLOTRAN is developed under an open source, GNU LGPL license. The term ‘lambda’ is used here because  $\lambda$  is a  
70 key parameter in the SXTM, which quantifies thermodynamic favorability of aerobic respiration of OM. The  
71 connection between the unique reaction network developed for each FTICR-MS sample hinges on the use of  
72 PFLOTRAN’s reaction sandbox capability (Hammond, 2022). The reaction sandbox gives the ability to define  
73 additional custom, kinetic reactions beyond standard formulations (e.g., mineral precipitation-dissolution, Michaelis-  
74 Menten, etc.). The Lambda-PFLOTRAN workflow enables upscaling by using molecular-scale information to inform  
75 larger scale biogeochemical processes occurring throughout a watershed which can be simulated with PFLOTRAN.  
76 Herein we describe the Lambda-PFLOTRAN workflow process including the governing expressions, workflow steps,  
77 data requirements, as well as the associated assumptions and limitations. Two illustrative test cases are also included  
78 to demonstrate the workflow.

## 79 **2 Methods**

### 80 **2.1 Conceptual Model**

81 Respiration modeling herein is based on thermodynamic theory by Desmond-Le Quemener and Bouchez (2014) which  
82 was updated for multiple OM formulas by Song et al. (2020). The generalized form of OM molecule is assumed to  
83 take the form of  $C_aH_bN_cO_dP_eS^z_f$ . Each molecular formula then undergoes respiration (i.e., reaction with oxygen) based  
84 on the following general reaction expression:



87 This generalized expression is used to describe the oxidation of any OM molecule,  $i$ , and has been normalized to one  
88 mole of biomass (BM) produced. BM is assumed to have a formula of  $CH_{1.8}O_{0.5}N_{0.2}$  (Stephanopoulos et al.,  
89 1998; Kleerebezem and Van Loosdrecht, 2010).  $OM_i$  represents the OM molecules as informed by FTICR-MS. Each  
90  $y$  represents the reaction stoichiometry for that reactant ( $y < 0$ ) or product ( $y > 0$ ). While this expression is specific for  
91 cases where oxygen is the electron acceptor, such an expression could be updated for alternative electron acceptors.

92 Substrate-explicit thermodynamic modeling expressions developed from Song et al. (2020) were implemented in a  
93 reaction sandbox within PFLOTRAN. The expressions were implemented in a general manner allowing for flexibility  
94 in handling variations in FTICR-MS data and several user adjustable analysis configurations.

95 The microbial growth kinetics are described by Eq. (2):

$$96 \quad \mu_i^{kin} = \mu^{max} \exp\left(-\frac{\alpha|y_{OM,i}|}{1000V_h[OM_i]}\right) \exp\left(-\frac{\alpha|y_{O_2,i}|}{1000V_h[O_2]}\right), \quad (2)$$

97 where  $\mu_i^{kin}$  is the unregulated uptake rate of reaction for  $OM_i$  [ $hr^{-1}$ ],  $\mu^{max}$  is the maximal microbial growth rate [ $hr^{-1}$ ],  $y_{OM,i}$  is the stoichiometry for  $OM_i$  [ $mol-OM \cdot mol-biomass^{-1}$ ],  $V_h$  is microbial harvest volume [ $m^3$ ]. Given the  
98 physical interpretation of  $V_h$  as the microbial harvest volume, it is assumed here that the value of  $V_h$  is the same for  
99 both  $OM_i$  and  $O_2$ ,  $[OM_i]$  is the organic matter concentration of  $OM_i$  [ $mol-OM \cdot L^{-1}$ ],  $y_{O_2,i}$  is the stoichiometry for  $O_2$   
100 for respiration of  $OM_i$  [ $mol-O_2 \cdot mol-biomass^{-1}$ ],  $[O_2]$  is oxygen concentration [ $mol-O_2 \cdot L^{-1}$ ],  $\alpha$  is a microbial unit  
101 conversion [ $mol-biomass$ ] and 1000 is the conversion of  $m^3$  to L.  
102

103 Further, using a cybernetic modeling approach (after Song et al., 2018), all the unregulated uptake rates ( $\mu_i^{kin}$ ) are  
104 normalized by the sum of unregulated uptake rates across all reactions,  $i$  following Eq. (3):

$$105 \quad u_i = \frac{\mu_i^{kin}}{\sum_{i=1}^n \mu_i^{kin}} \quad (3)$$

106 where  $u_i$  is the fraction of the unregulated rate [-]. The final regulated rate,  $r_i$  [ $hr^{-1}$ ] for each reaction is then computed  
107 following Eq. (4):

$$108 \quad r_i = u_i \mu_i^{kin}, \quad (4)$$

109 For implementation within PFLOTRAN, the use of inhibition terms was required to prevent negative concentrations  
110 once a reactant is nearly depleted. For a reaction to proceed, all reactant species must be present above a minimum  
111 concentration even if the molecules do not explicitly control the respiration rate (i.e., species other than OM and  $O_2$ ,  
112 Eq. (2). If a reactant concentration falls below a threshold concentration, the respiration rate is inhibited. Reactant  
113 inhibition is computed by Eq. 5 (Kinzelbach et al., 1991) for reactant species  $j$ :

$$114 \quad I_j = 0.5 + \frac{\arctan([C_j] - C_{thj}) \cdot f}{\pi}, \quad (5)$$

115 where  $C_{th,i}$  is the threshold concentration [M],  $f$  is the threshold scaling factor [-]. The default  $C_{thj}$  is  $10^{-20}$  M.

116 The reaction rates are also inhibited by the microbial carrying capacity of the system,  $I_{cc}$ , as follows in Eq. (6):

$$117 \quad I_{cc} = 1 - \frac{[BM]}{CC} \quad (6)$$

118 where [BM] is the biomass concentration [ $mol-BM \cdot L^{-1}$ ], CC is the biomass carrying capacity [ $mol-BM \cdot L^{-1}$ ].  $I_{cc}$  has a  
119 non-negativity constraint, so if  $[BM] > CC$ , then  $I_{cc} = 0$ .

120 These inhibition factors are applied to the overall rate expression as shown in Eq. (7).

121  $r_{i,inhibited} = r_i I_{CC} \prod I_j \quad \forall y_{i,j} < 0,$  (7)

122 The overall individual species rates,  $d[C_j]/dt$ , [mol-species·L<sup>-1</sup>·hr<sup>-1</sup>] are then computed as follows with Eq. (8):

123  $\frac{dC_j}{dt} = (\sum_{i=1}^n y_{i,j} r_{i,inhibited}) [BM],$  (8)

124 where  $j$  is the species index. The total number of species includes 7 general species (i.e., HCO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup>, HPO<sub>4</sub><sup>-</sup>, HS<sup>-</sup>,  
 125 H<sup>+</sup>, O<sub>2</sub>, BM (i.e., Eq (1)) and the OM species considered (i.e., typically 10).  $i$  is the reaction index,  $n$  is total number  
 126 of reactions as based on the total number of OM species (typically, with this workflow  $n = 10$ ).  $y_{i,j}$  is the coefficient  
 127 for species  $j$  in reaction  $i$ .

128  
 129 The expression for biomass is also modified to account for biomass decay (note all biomass stoichiometries are 1 by  
 130 definition):

131  $\frac{dBM}{dt} = (\sum_{i=1}^n y_{i,j} r_{i,inhibited}) [BM] - k_{deg} [BM],$  (9)

132 where  $k_{deg}$  is the biomass decay rate [hr<sup>-1</sup>].

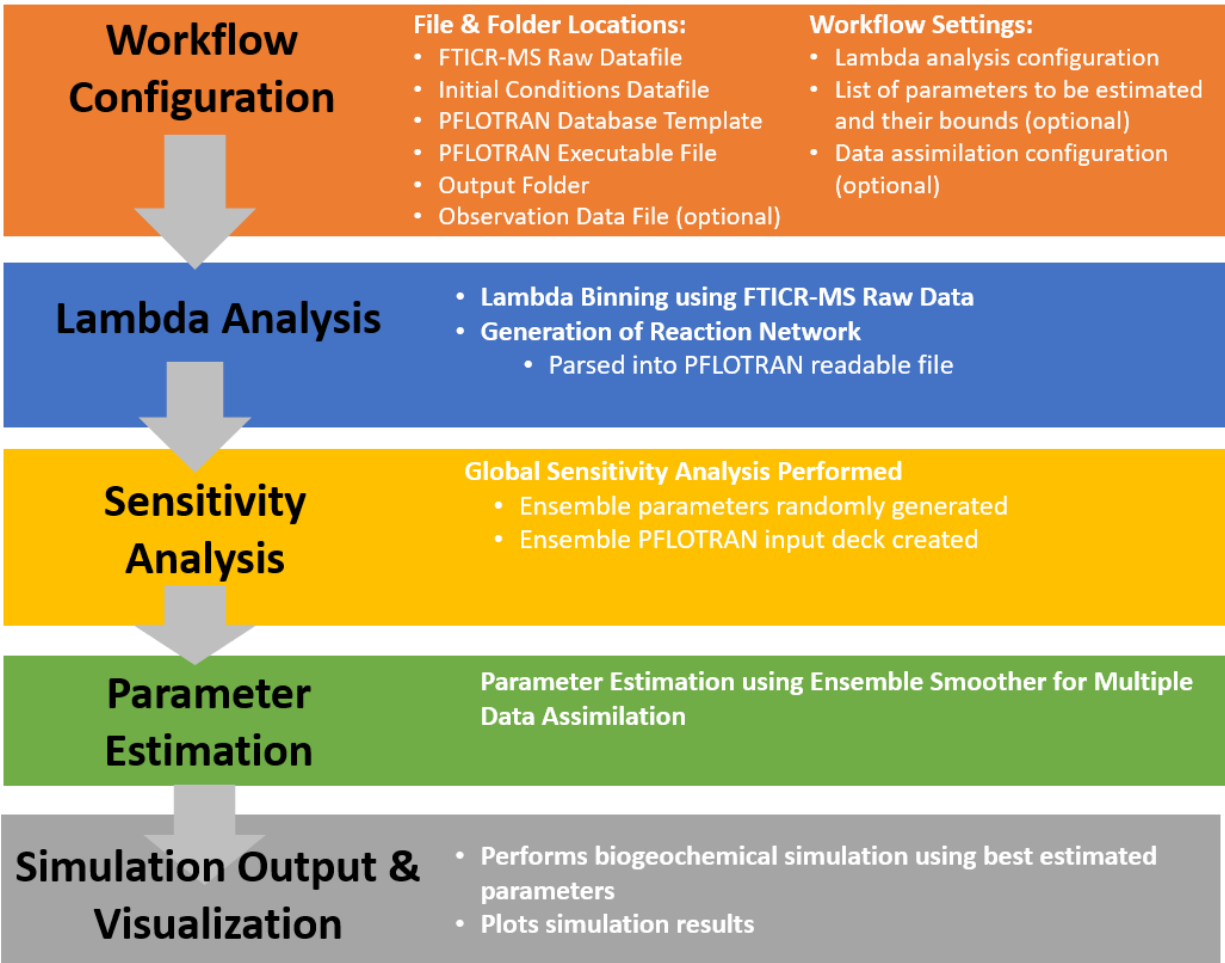
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## 134 2.2 Lambda Analysis and Binning

135 To reduce the number of organic compounds considered in the simulation, OM molecules are grouped, or binned,  
 136 based on their  $\lambda$  value computed by Eq. (10):

137  $\lambda = \frac{\Delta G_{r,anabolic} + \Delta G_{r,dissipation}}{(-\Delta G_{r,catabolic})},$  (10)

138 where  $\Delta G$  are the Gibbs energies for the anabolic and catabolic reactions and the associated dissipation energy,  
 139 respectively. The value of  $\lambda$  is indicative of how many times the catabolic reaction needs to be completed to provide  
 140 the energy required to synthesis one mole of biomass. Lower  $\lambda$  values suggest higher thermodynamic favorability of  
 141 OM respiration. Using the chemical formula determined for each OM molecule, the energy balance equations are  
 142 solved providing the overall reaction stoichiometry Eq. (1) and the  $\lambda$  is calculated. Using the  $\lambda$  value for each molecule,  
 143 the cumulative probability distribution for the sample is produced (Figure 2).

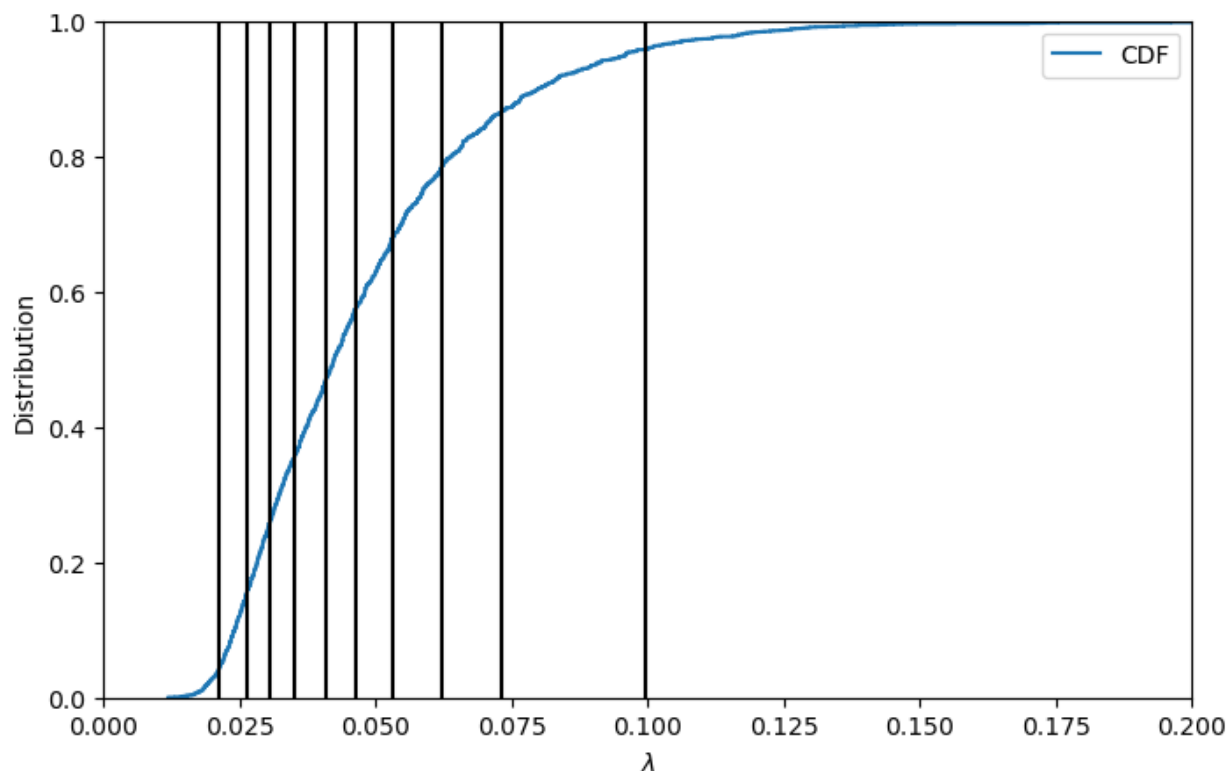


144

145 **Figure 1:** Flow Chart of the Lambda-PFLOTRAN Workflow.

146

147 It is this conversion from individual compounds to a distribution that is critical for reducing the entire sample down  
 148 to a representative set of expressions. The  $\lambda$  bins are then formed by splitting the cumulative probability distribution  
 149 into equally weighted sections as which to define the overall sample by. The illustrative example shown in Fig. 2  
 150 demonstrates the sample distribution being divided into 10 sections (i.e., in this case each section contains 10% of the  
 151 overall sample distribution).



152  
 153 **Figure 2:** Lambda binning to convert raw FTICR-MS into a representative reaction network using the cumulative probability  
 154 distribution function (CDF) for Test Case 1a. Vertical lines display the average  $\lambda$  value for each of the 10 bins (left to right,  $\lambda$  bin  
 155 1 to 10).

156  
 157 Each section is used to determine a representative organic matter formula and the associated reaction and  
 158 stoichiometry of that  $\lambda$  bin. The group of representative reactions (one per bin) is called the reaction network. A  
 159 demonstrative reaction network defined by  $\lambda$  analysis and binning is shown in Table 1.

160  
 161 **Table 1:** Reaction Network Developed from Lambda Theory for Test Case 1a

Bin Number	Representative Organic Matter Species Formula	$\lambda$	$y_{OM}$	$y_{HCO_3^-}$	$y_{NH_4^+}$	$y_{HPO_4^{2-}}$	$y_{HS^-}$	$y_{H^+}$	$y_{O_2}$
1	$C_{31}H_{44}N_{0.33}O_{4.8}P_{0.6}S_{0.3}$	0.021	-0.05	0.64	-0.17	-0.18	0.03	0.02	-1.07
2	$C_{26}H_{39}N_{0.20}O_{7.0}P_{0.6}S_{0.1}$	0.026	-0.07	0.68	-0.10	-0.19	0.04	0.01	-1.06
3	$C_{22}H_{36}N_{0.24}O_{7.5}P_{0.5}S_{0.1}$	0.031	-0.08	0.69	-0.02	-0.18	0.04	0.01	-1.06
4	$C_{20}H_{32}N_{0.28}O_{7.3}P_{0.4}S_{0.1}$	0.035	-0.08	0.72	-0.08	-0.18	0.04	0.01	-1.05
5	$C_{19}H_{29}N_{0.48}O_{7.9}P_{0.3}S_{0.2}$	0.041	-0.09	0.79	-0.17	-0.16	0.03	0.02	-1.04
6	$C_{18}H_{26}N_{0.68}O_{8.1}P_{0.2}S_{0.2}$	0.046	-0.10	0.85	-0.27	-0.13	0.02	0.02	-1.03
7	$C_{17}H_{24}N_{0.69}O_{8.1}P_{0.2}S_{0.2}$	0.053	-0.11	0.90	-0.32	-0.12	0.02	0.02	-1.02
8	$C_{15}H_{20}N_{0.67}O_{7.6}P_{0.2}S_{0.2}$	0.062	-0.13	0.94	-0.42	-0.11	0.02	0.03	-1.00

9	$C_{13}H_{19}N_{1.13}O_{87.4}P_{0.1}S_{0.2}$	0.073	-0.15	1.01	-0.48	-0.03	0.01	0.03	-1.00
10	$C_{10}H_{15}N_{1.56}O_{6.5}P_{0.1}S_{0.2}$	0.100	-0.21	1.17	-0.75	0.12	0.01	0.04	-0.97

162  
163 Currently, the representative OM molecule that defines each bin is computed as the average chemical formula of all  
164 the molecules present in that  $\lambda$  section. The disadvantage of this approach is that unrealistic compounds are defined  
165 as representative molecules instead of realistic molecules. The issue with selecting a single, but real compound, from  
166 within each  $\lambda$  section resides in chemical complexity and variation - for instance some molecules may contain low  
167 levels of phosphorous or sulfur and others may not contain either element in the chemical formula. Thus, requiring  
168 the representative chemical formula to be a real compound present in the sample would create basis which would  
169 propagate through the reaction network and into the resulting biogeochemical simulation results.

## 170 **2.3 Lambda-PFLOTRAN Workflow**

171 The Lambda-PFLOTRAN workflow digests raw FTICR-MS data, calculates the  $\lambda$  distribution for the sample,  
172 generates the  $\lambda$  bins and corresponding reaction network, and completes a biogeochemical simulation using  
173 PFLOTRAN. Further, we incorporated sensitivity analysis and ensemble data assimilation to enable an in-depth  
174 exploration of the impact of reaction parameters on respiration as well as a straightforward parameter estimation  
175 method to fit model parameters to experimental data.

176 The workflow is implemented through a user-friendly Jupyter notebook interface (Kluyver et al., 2016) where a user  
177 can configure the simulation parameters by adjusting initial concentrations,  $\lambda$  binning configuration, parameter values  
178 and/or ranges, and data assimilation options. Based on the user's data file and the associated parameters, scripts within  
179 the Jupyter notebook write the corresponding PFLOTRAN input files, including OM molecules and aqueous  
180 chemistry. The PFLOTRAN simulations are completed locally through a Docker container making this capability  
181 much more user-friendly and accessible. The progress of the data assimilation tool used for parameter fitting is  
182 illustrated within the Jupyter notebook. The resulting best fit final biogeochemical simulation is output visually with  
183 plots and as a text file (when applicable).

184 The Lambda-PFLOTRAN workflow steps are shown in Figure 1 and described in detail in the following subsections:

### 185 **2.3.1 Step 1 – Workflow configuration**

186 The first step is to set up the workflow configuration for a Lambda-PFLTORAN application. This includes specifying  
187 the file and folder locations of the following information: 1) FTICR-MS raw data file (.csv), 2) initial species  
188 concentrations file (.csv) that includes starting molar concentrations for  $HCO_3^-$ ,  $NH_4^+$ ,  $HPO_4^{2-}$ ,  $HS^-$ ,  $H^+$ ,  $O_2$  (aq), BM  
189 and total organic carbon (TOC), 3) PFLOTRAN database template file, 4) PFLOTRAN executable file, 5) workflow  
190 output folder, and if completing parameter estimation, (6) the data observation file (.csv), if applicable.

191 The user is also asked to configure workflow settings related to: (1) the lambda analysis configuration, including  
192 number of  $\lambda$  bins and method to define the  $\lambda$  bins (i.e., cumulative vs uniform); (2) the respiration modeling parameter



193 setup, including the list of the parameters to be estimated and their associated upper and lower bounds and (3) the data  
194 assimilation configuration (see below).

### 195 **2.3.2 Step 2 – Organic Matter Chemistry using Lambda Analysis**

196 With only an input of FTICR-MS data, the workflow first performs the lambda analysis (Section 2.2) to group OM  
197 molecules into various  $\lambda$  bins based on each compound's thermodynamics (Figure 2) and produce the corresponding  
198 reaction network for respiration (Table 1). The default number of  $\lambda$  bins is 10, although this can be adjusted in the  
199 workflow configuration by the user, if desired. The generated reaction network is then automatically parsed by the  
200 workflow into a text file that can be read by PFLOTRAN.

### 201 **2.3.3 Step 3 – Sensitivity Analysis using Mutual Information**

202 This step performs the global sensitivity analysis on the parameters to be estimated. Ensemble parameters are first  
203 generated by randomly sampling from their predefined ranges in the configuration step and saved into an HDF5 file.  
204 Then, the workflow generates a PFLOTRAN input deck to conduct ensemble simulations using the ensemble  
205 parameters. The generated ensemble model states enables a global sensitivity analysis using mutual information  
206 (Cover and Thomas, 2006; Jiang et al, 2022) as follows:

$$207 \quad I(X;Y) = H(Y) - H(Y|X) = \sum_{x=x} \sum_{y=y} p(x,y) \log \left( \frac{p(x,y)}{p(x)p(y)} \right), \quad (11)$$

208 where  $x$  and  $y$  are the specific values of  $X$  and  $Y$ , respectively;  $H(Y)$  is the Shannon's entropy of  $Y$ ;  $H(Y|X)$  is the  
209 conditional entropy of  $Y$  given  $X$ ;  $p$  is the probability density function. Higher  $I$  indicate stronger sensitivity between  
210  $X$  and  $Y$ . Besides sensitivity analysis, the ensemble parameter/states also serve as the prior information for parameter  
211 estimation at the next step.

### 212 **2.3.4 Step 4 – Parameter Estimation using Ensemble Smoother for Multiple Data Assimilation**

213 The workflow adopts Ensemble Smoother for Multiple Data Assimilation (Emerick and Reynolds, 2013; Jiang et al,  
214 2021), abbreviated as ESMDA, for data assimilation in this step. Rooted in ensemble Kalman filter, ESMDA is an  
215 iterative data assimilation approach that assimilates the observations on the entire time period for multiple times to  
216 reduce the uncertainty of the estimated or posterior parameters. During each iteration of ESMDA, the model  
217 parameters are updated based on the following equation:

$$218 \quad m_{k,l}^u = m_{k,l}^f + C_{MD,l}^f (C_{DD,l}^f + \alpha_l C_D)^{-1} \left( d_{obs} + \sqrt{\alpha_l} C_D^{\frac{1}{2}} z_k - d_{k,l}^f \right), \quad k = 1, \dots, N_e \text{ and } l = 1, \dots, L, \quad (12)$$

219 where the subscripts  $k$  and  $l$  are the indices of the ensemble member and the iteration, respectively; the superscripts  $u$   
220 and  $f$  are the updated and forecast parameters or states, respectively;  $N_e$  is the number of ensemble members;  $L$  is the  
221 number of iterations;  $m_{k,l}^f$  and  $m_{k,l}^u$  are the  $k$ th ensemble member of the forecast/prior and updated/posterior  
222 parameters, respectively, at the  $l$ th iteration;  $d_{obs}$  is the observation;  $z_k$  is the observation noise sampled from

223 independent standard normal distributions for the  $k$ th ensemble member;  $d_{k,l}^f$  is the  $k$ th ensemble member of the  
224 predicted observation states by the model using  $m_{k,l}^f$ ;  $C_{MD,l}^f$  is the cross-covariance matrix between the prior parameters  
225  $m_l^f$  and the predicted observation states  $d_l^f$ ;  $C_{DD,l}^f$  is the auto-covariance matrix of the predicted observation states  $d_l^f$ ;  
226  $C_D$  is the auto-covariance matrix of the observation error; and  $\alpha_l$  is the inflation coefficient at the  $l$ th iteration with the  
227 sum of all  $\alpha_l$  equal to one.

228  
229 Here, the assimilation starts with taking the ensemble model parameters/states in Step 3 and the provided observations,  
230 and calculates the posterior parameters using ensemble Kalman filter, updates the prior parameters with the current  
231 posterior for the next iteration, and then repeats the whole process for multiple times (typically 3 to 5 iterations, as  
232 defined by the user). The final estimated parameters are obtained from the posterior parameter at the last iteration and  
233 are updated in the parameter HDF5 file. The parameter estimation is implemented in a way that allows assimilating  
234 either a single (e.g., Test Case 1) or multiple observed species simultaneously through a simple change of the inputs.  
235 For example, if temporal experimental or field data is available for oxygen, pH, and total carbon, all these data sources  
236 could be simultaneously fit to, with only minor adjustments to Jupyter notebook.

### 237 **2.3.5 Step 5 – Simulation Output and Visualization**

238 The last step performs the ensemble simulation of the biogeochemical modeling a final time using the estimated  
239 parameters in Step 4. Optionally, users can further pick the realization with the best performance. The user has the  
240 option to select their preferred goodness of fit metric from the following options as a means for selecting the best  
241 performing simulation: R-squared ( $R^2$ ), Root Mean Squared Error (RMSE), Modified Kling-Gupta Efficiency  
242 (mKGE), Nash-Sutcliffe Model Efficiency Coefficient (NSE), or Correlation Coefficient (CorC). Based on the  
243 selection, the final time series of aqueous chemistry, oxygen consumption,  $\text{CO}_2$  production, and lambda binned, and  
244 total organic carbon concentrations will be computed and plotted.

## 245 **3 Test Cases**

### 246 **3.1 Test Case 1 - Oxygen Depletion Incubation Experiments.**

247 In the first illustrative example, the workflow was used to fit  $\mu_{max}$  to laboratory incubation experiments where oxygen  
248 levels were measured over two hours in a closed reactor. The incubation experiments were completed as part of the  
249 Worldwide Hydrobiogeochemistry Observation Network for Dynamic River Systems (WHONDRS) program  
250 (Goldman et al, 2020). For these incubations, sediment was taken from three locations within a stream (i.e., upstream  
251 [Test Case 1a], midstream [Test Case 1b], and downstream [Test Case 1c]) in the Yakima River Basin in Washington,  
252 USA for subsequent laboratory respiration experiments. FTICR-MS was used to determine the OM chemistry from  
253 each sediment sample, resulting in variable formulae being identified in each sample. Formula assignments for all the  
254 samples included herein were completed using formultitude (Tolic et al., 2017). Total dissolved organic carbon  
255 concentration paired with the FTICR-MS sample and biomass measurements taken at the start of each experiment  
256 were used as the initial concentrations for each of the simulations. Due to the absence of quantitative data related to

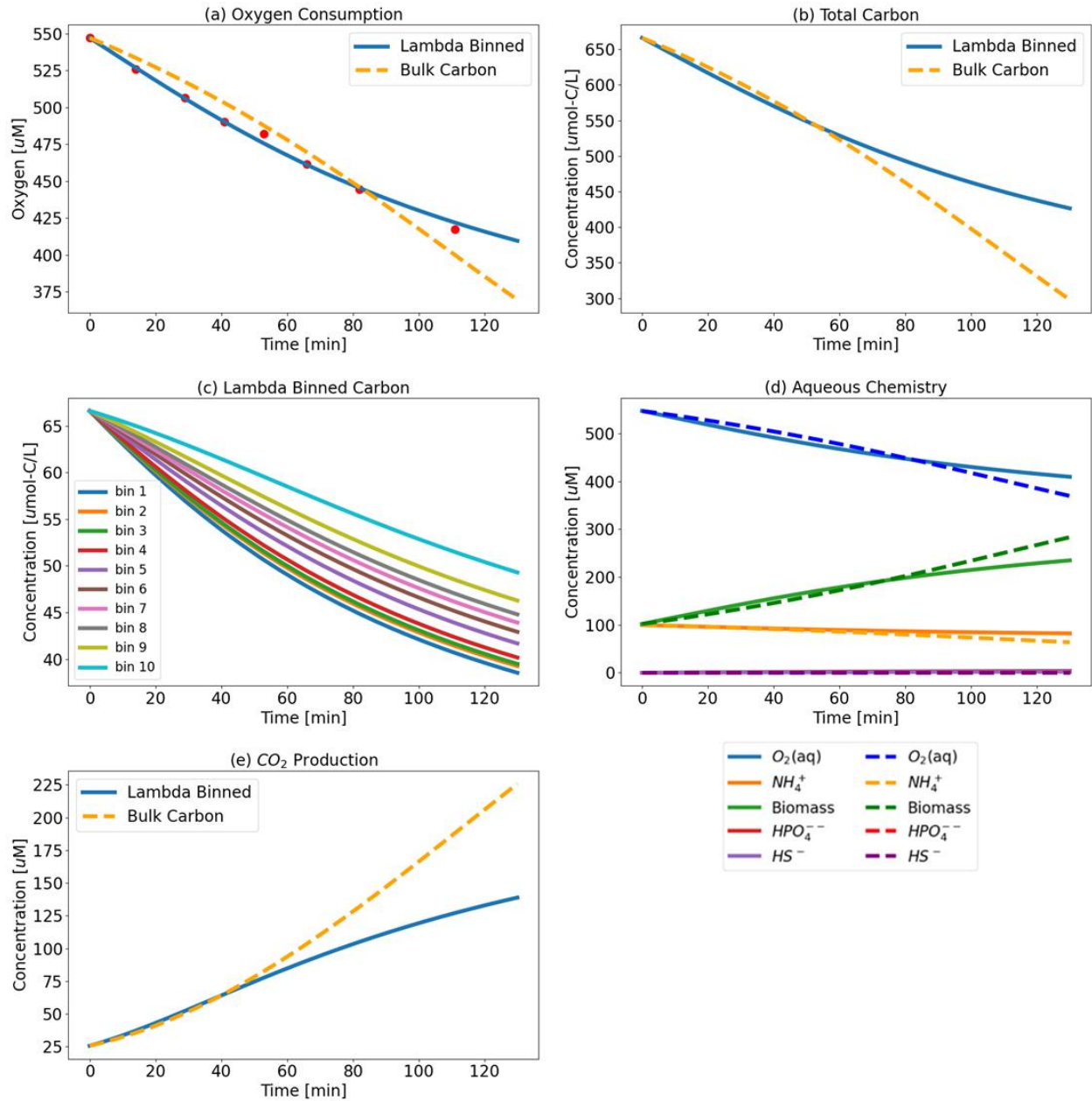
257 how the total carbon mass is distributed between various the OM compounds, the total carbon concentration (on a per-  
258 C basis) was assumed to be split equally between each of the  $\lambda$  bins. The total organic carbon concentration was  
259 distributed into each  $\lambda$  bin using Eq. (13). While this assumption results in equal distribution of carbon between the  
260 bins, consequently, it assigns different initial species concentrations due to varying carbon concentrations between the  
261 molecules.

$$262 \quad [C_{\lambda bin}]_0 = \frac{[TOC]}{n_{\lambda bin} n_{C_{\lambda bin}}} \quad (13)$$

263 Where:  $[C_{\lambda bin}]_0$  is the initial species concentration in each  $\lambda$  bin [ $\text{mol} \cdot \text{L}^{-1}$ ];  $TOC$  is the total organic carbon measured  
264 [ $\text{mol-carbon} \cdot \text{L}^{-1}$ ];  $n_{\lambda bin}$  is the number of  $\lambda$  bins [-]; and  $n_{C_{\lambda bin}}$  is the number of carbon molecules in the assumed  
265 formula for the  $\lambda$  bins [ $\text{mol-carbon} \cdot \text{mol-molecule}^{-1}$ ].

266 Using the Lambda-PFLOTRAN workflow, the FTICR-MS data from each laboratory experiment was digested into  
267 the corresponding  $\lambda$  bins to create the individual reaction network. The Jupyter Notebook for this example is  
268 “Test\_Case1-WHONDRS.ipynb” and is available at <https://doi.org/10.15485/2281403>.

269  $\mu_{max}$  was fit to the provided experimental oxygen data. The final lambda binned fit, along with corresponding carbon  
270 consumption (individual and total) and aqueous chemistry is displayed in Figure 3 (and in the supporting information  
271 Fig. S1 and S2 for Test Cases 1b [midstream] and 1c [downstream], respectively). To evaluate the use of lambda  
272 binned OM obtained from FTICR-MS (Figure 3), the workflow was also run for a baseline case where  $\mu_{max}$  was fit  
273 again, but this time assuming a generic bulk OM form of  $\text{CH}_2\text{O}$  for comparison. Fitted  $\mu_{max}$  values for the lambda  
274 binned model is  $0.25 \text{ min}^{-1}$  ( $R^2 = 0.99$ ) and fitted  $\mu_{max}$  to the bulk OM  $\text{CH}_2\text{O}$  model is  $0.032 \text{ min}^{-1}$  ( $R^2 = 0.96$ ).  $V_h$  and  
275  $CC$  are fixed at assumed values of  $10 \text{ m}^3$  and  $1 \text{ M}$ , respectively in both simulations.



276  
 277 **Figure 3:** Test Case 1a Results – (a) Oxygen Consumption where Lambda-PFLOTTRAN workflow was used to fit (blue line) to  
 278 experimental respiration data (red dots) and (b) Total Carbon Consumption; (c) Individual Organic Matter Consumption by  $\lambda$  bin;  
 279 and (d) biogeochemistry including  $O_2$  (aq) (blue); Biomass (green);  $NH_4^+$  (orange);  $HS^-$  (purple); and  $HPO_4^-$  (red); and (e)  $CO_2$   
 280 production for the upstream incubation. The dashed orange lines (in a, b and e) show simulation results assuming a generic OM  
 281 species of  $CH_2O$  for comparison.

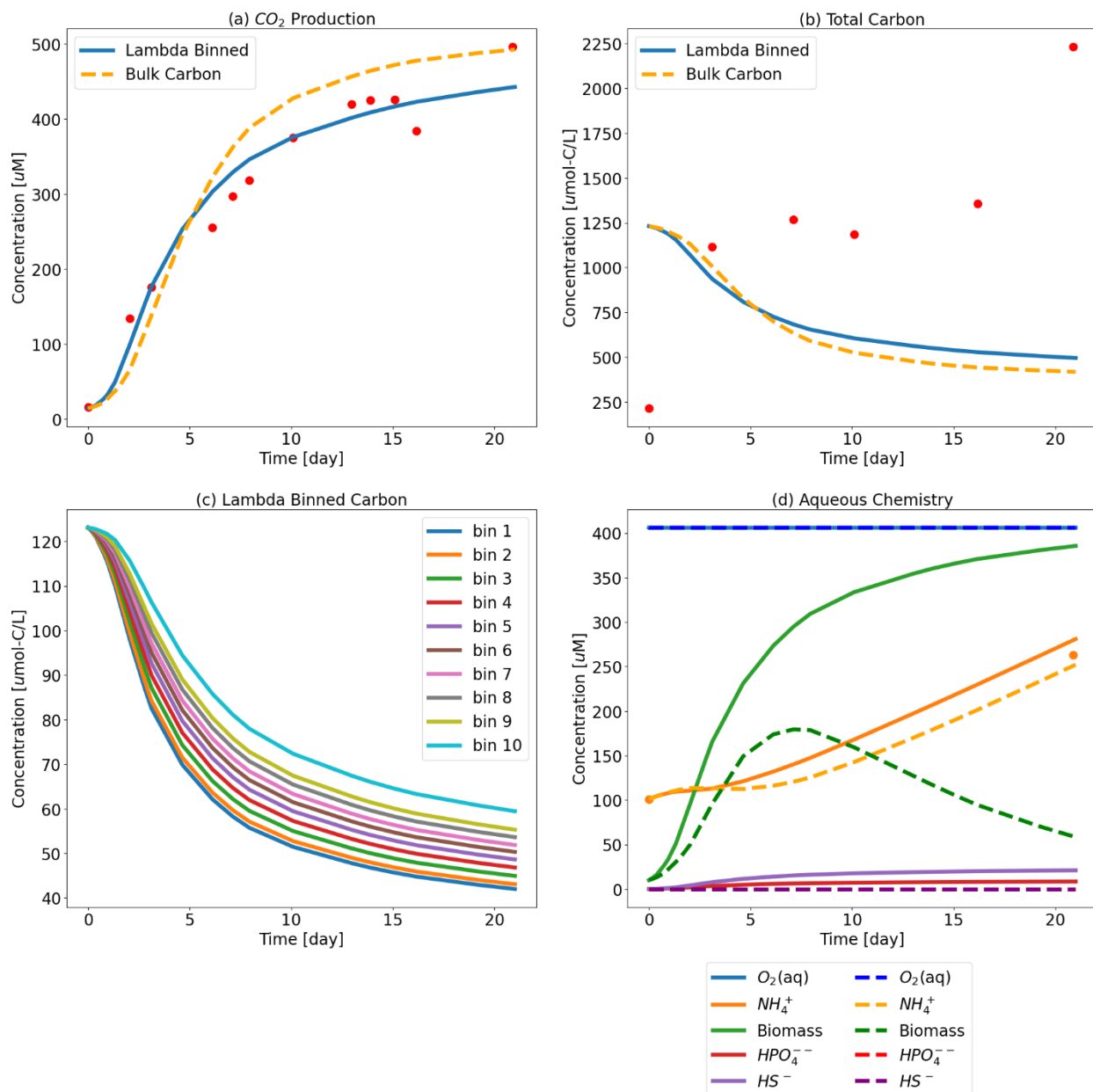
282 However, even over the short time frame of this simulation (i.e., only 120 minutes), the difference between assuming  
 283 the generic  $CH_2O$  and using the more detailed organic matter chemistry resulted in different predictions of total carbon  
 284 and  $CO_2$  generation. The bulk OM model predicts more carbon consumption and greater  $CO_2$  production than the  
 285 lambda binned model. The bulk OM model estimates that 50% of the initial total carbon is consumed over the first

286 120 mins, whereas the lambda binned model predicts 34% consumption. Similarly, the bulk OM model estimates  
287 approximately 35% more CO<sub>2</sub> generation as compared the lambda binned model. The effects on aqueous chemistry  
288 over this short duration are more muted, albeit still present.

### 289 **3.2 Test Case 2 - Respiration Incubation Experiments.**

290 Test Case 2 uses soil respiration incubation data from Ward et al. (2023) aimed at investigating the influence of soil  
291 type, oxygen condition (aerobic vs. anaerobic), and seawater exposure (fresh vs. saline) on respiration extent and rate.  
292 For these experiments, temporal measurements were collected for CO<sub>2</sub> generation, dissolved organic carbon (DOC),  
293 organic matter formulas via FTICR-MS and other bulk aqueous chemistry (i.e., pH, NH<sub>4</sub><sup>+</sup>, and other metals and ions)  
294 creating a rich dataset for calibration of system specific lambda model parameters. These incubations were setup by  
295 adding dry soil to the reactor and then adding water (resulting in a soil:water ratio ranging from 1:11 to 1:16). The soil  
296 and water were shaken vigorously for five minutes, and then sampled for the initial time point prior to officially  
297 starting the incubation. For the aerobic experiments, the reactor headspace was cycled every 24 hours to measure CO<sub>2</sub>  
298 generated but also to ensure the system was kept aerobic; this was only performed five days per week, with no  
299 measurements taken on the weekend due to logistical constraints. Upon experiment completion, the increase in DOC  
300 concentrations indicated organic carbon was being kinetically released from the soil into the aqueous phase over the  
301 course of the 21-day experiment. Similarly, measured NH<sub>4</sub><sup>+</sup> concentrations also increased during the experiment. To  
302 address this within our reactive transport model, a source of nitrogen was assumed to be released from the soil as well  
303 ( $N_{release}$ ). Both carbon and nitrogen release are included in this example and are assumed to follow a zero-order constant  
304 release rate. Any organic carbon released from the soil was fractionated into each  $\lambda$  bin on the same per-carbon basis  
305 assumed for the initial total organic carbon. This was implemented through a dependent function that calculated the  
306 release of carbon into each  $\lambda$  bin based on a fitted single bulk  $k_{release}$  rate. Mathematically in PFLOTTRAN the constant  
307 oxygen conditions were implemented through a gas-liquid partitioning expression with a fast exchange term. These  
308 three additional processes were added to describe the experimental conditions of Test Case 2 more accurately (i.e.,  
309 release of carbon, nitrogen and sustained aerobic conditions); however, a PFLOTTRAN input deck can be expanded  
310 and customized to include a host of additional processes and full geochemistry for a specific system of interest. For  
311 instance, aqueous complexation, mineral dissolution and precipitation, sorption, and redox reactions can be added, all  
312 of which can influence the resultant pH and carbon, nitrogen, and other nutrient dynamics.

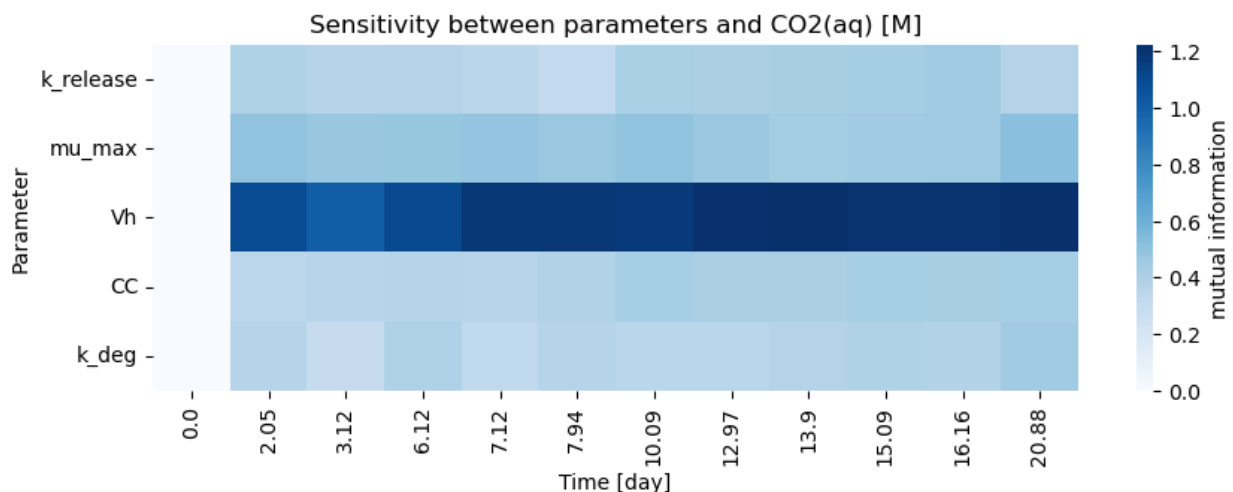
313 The workflow was used to fit  $\mu_{max}$ ,  $V_h$ ,  $CC$ ,  $k_{deg}$ , as well as  $k_{release}$ , to the temporal CO<sub>2</sub> generation for a single aerobic  
314 soil incubation (Figure 5). The Jupyter Notebook for this example is “Test\_Case2-Colloids.ipynb”.



316  
 317 **Figure 4.** Test Case 2 Results – (a) CO<sub>2</sub> production where Lambda-PFLOTTRAN workflow was used to fit (blue line) to  
 318 experimental respiration data (red dots) and (b) the corresponding Total Organic Carbon; (c) Individual Organic Matter  
 319 Consumption by  $\lambda$  bin, and (d) the corresponding biogeochemistry including O<sub>2</sub> (aq) (blue); Biomass (green); NH<sub>4</sub><sup>+</sup> (orange); HS<sup>-</sup>  
 320 (purple); and HPO<sub>4</sub><sup>-</sup> (red). Dots indicate experimental data. The dashed orange lines in the top two figures show simulation results  
 321 assuming a generic OM species of CH<sub>2</sub>O for comparison. Fitted parameters for lambda binned model are  $k_{\text{release}} = 5.5 \times 10^{-12} \text{ day}^{-1}$ ;  
 322  $\mu_{\text{max}} = 37.6 \text{ day}^{-1}$ ,  $V_h = 5.0 \text{ m}^3$ ,  $CC = 0.12 \text{ M}$ , and  $k_{\text{deg}} = 1 \times 10^{-3} \text{ day}^{-1}$  ( $R^2 = 0.953$ ) and fitted bulk OM CH<sub>2</sub>O model values are  $k_{\text{release}}$   
 323  $= 2.0 \times 10^{-12} \text{ day}^{-1}$ ;  $\mu_{\text{max}} = 47 \text{ day}^{-1}$ ,  $V_h = 1.0 \text{ m}^3$ ,  $CC = 0.77 \text{ M}$ , and  $k_{\text{deg}} = 0.15 \text{ day}^{-1}$  ( $R^2 = 0.909$ ).

324 For the purposes for showcasing the workflow, five parameters were estimated in this test case example, and as a  
 325 result the models are over parametrized given the amount of data available. Parameter sensitivity over the course of

326 simulation time is shown in Figure 5 and suggests this system is highly sensitive to  $V_h$ . It should be noted that both  
 327 these model fits are also highly sensitive to the allowable parameter space as user defined by the lower and upper  
 328 parameter bounds.

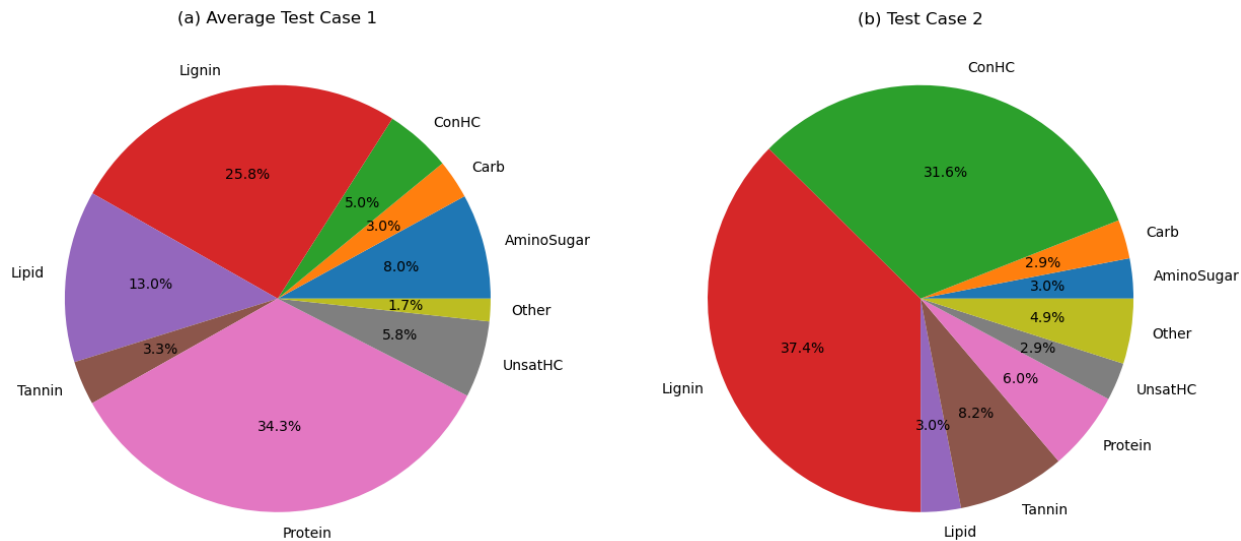


329  
 330 **Figure 5.** Test Case 2 - Sensitivity Analysis Output during Parameter Estimation. The sensitivity of five fitted parameters ( $k_{\text{release}}$ ,  
 331  $\mu_{\text{max}}$ ,  $V_h$ , CC, and  $k_{\text{deg}}$ ) on temporal aqueous  $\text{CO}_2$  concentrations as a function of time.

332 Any additional experimental data, either collected during incubations or through independent experiments (e.g.,  
 333 carbon release from the soil in an abiotic system), would be expected to help constraint the model and improve  
 334 parameterization. Additionally, it is unclear why the model is unable to capture the total organic carbon behavior in  
 335 Test Case 2. One potential explanation is that some of the released organic carbon may not be fully bioavailable and  
 336 thus the model may be compensating for this by artificially reducing the concentration of OM available for respiration.

#### 337 4 Variability and Impact of Organic Matter Speciation

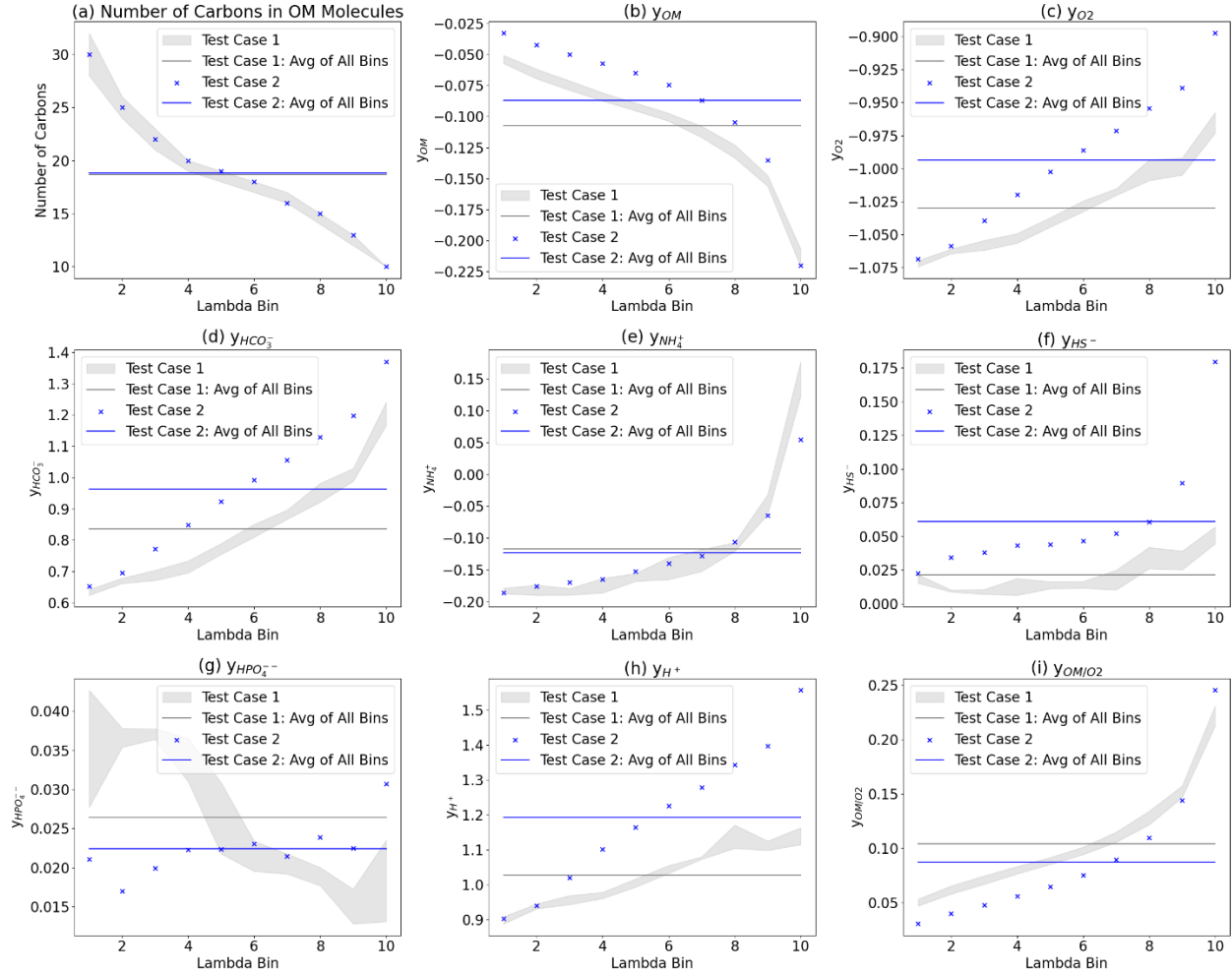
338 The variability in OM speciation was briefly assessed by comparing FTICR-MS data from Test Cases 1 and 2. Each  
 339 identified OM species was classified into one of nine compound classes. For Test Case 1, the average of the three Test  
 340 Case 1 samples (1a - upstream, 1b - midstream, and 1c - downstream) was computed. The predominant classes were  
 341 proteins ( $34 \pm 1\%$ ), lignin ( $26 \pm 1\%$ ), and lipids ( $13 \pm 2\%$ ), with the errors representing the standard deviation among  
 342 the Test Case 1a-c samples. The low standard deviation suggests consistent reproducibility in OM speciation for  
 343 samples taken from nearby locations. In contrast, OM in Test Case 2 was primarily composed of lignin (37.4%) and  
 344 concentrated hydrocarbons (32%). The full distribution of compound classes is presented in Figure 5.



345  
 346 **Figure 5.** Distribution of Organic Matter Compound Classes: (a) Test Case 1 and (b) Test Case 2.  
 347 Note: Test Case 1 is the average of Test Case samples 1a-c. ConHC = Condensed Hydrocarbon; UnsathHC = Unsaturated  
 348 Hydrocarbon

349  
 350 The influence of the sample OM speciation on the  $\lambda$  binned reaction networks was also assessed. Figure 6 illustrates  
 351 the impact of OM speciation on the corresponding  $\lambda$  binned reaction networks, with three key observations. First, the  
 352 variability in OM speciation between different samples is evident when comparing Test Case 1 and Test Case 2. To  
 353 enhance visual clarity, the range of Test Case 1 samples (1a-c) is depicted as a grey shaded region, showing the spread  
 354 between the minimum and maximum values of the three samples. For Test Case 2, data from the single FTICR-MS  
 355 sample is represented by blue dots. Test Case 1 and 2 have distinct  $\lambda$  derived reaction networks as indicated by the  
 356 little overlap between the grey region and the blue dots in Figures 6b-i.



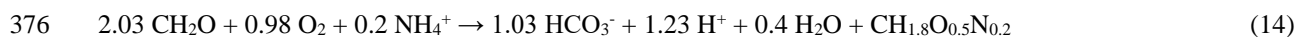


357  
 358 **Figure 6.** Comparison of Lambda-Binned Reaction Network Parameters: (a) number of carbons in the OM; stoichiometric  
 359 coefficient,  $y$ , for (b) OM, (c)  $O_2$ , (d)  $HCO_3^-$ , (e)  $NH_4^+$ , (f)  $HS^-$ , (g)  $HPO_4^-$ , (h)  $H^+$ ; and (i) ratio of OM/ $O_2$  coefficients for Test  
 360 Case 1a-c (grey dots); the average of all  $\lambda$  bins for Test Case 1 (grey line); Test Case 2 (blue x); and the average of all  $\lambda$  bins for  
 361 Test Case 2 (blue line). The grey shaded area highlights the range of values for Test Case 1a-c for better visual comparison.

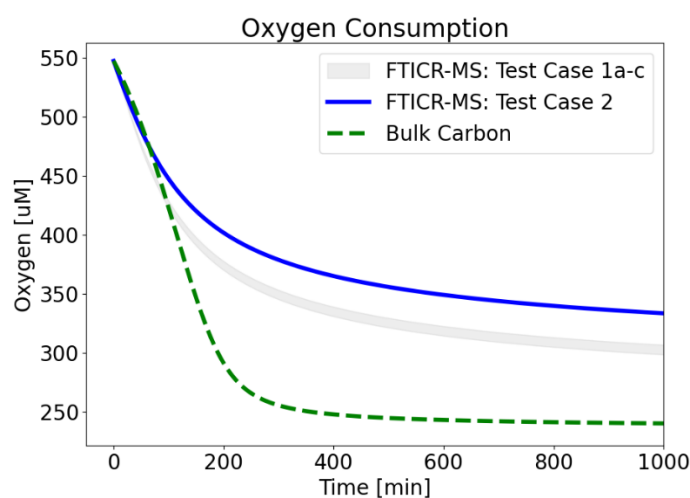
362  
 363 Second, the  $\lambda$  binning process captures the OM speciation variation within a sample. To illustrate this intrasample  
 364 variability, a line representing the average of all  $\lambda$  bins is shown on Figure 6 (grey line for Test Case 1, blue line for  
 365 Test Case 2). The difference between the reaction network coefficients (vertical axis) for the  $\lambda$  binning (grey shaded  
 366 area and blue dots) and the Test Case average lines highlights the extent of this variability. Finally, although the  $\lambda$   
 367 binning process resulted in a similar number of carbon atoms to OM molecules within each  $\lambda$  bin for both test cases  
 368 (Figure 6a), the resulting stoichiometric coefficients in the reaction networks differ significantly (Figures 6b-h). These  
 369 stoichiometric differences lead to variations in biogeochemical outcomes, such as OM-to-oxygen utilization ratios  
 370 during aerobic respiration (Figure 6i). These differences are due to the additional elements beyond carbon in the OM  
 371 molecules (i.e., nitrogen, oxygen, sulfur, hydrogen, and phosphorus).

372

373 Another important aspect is the comparison against assuming a generic, bulk OM composition, which does not account  
374 for OM speciation as informed by FTICR-MS or similar methods. The reaction network developed for a generic OM  
375 molecule of CH<sub>2</sub>O is shown in Eq. 14.



377 This reaction network is used in the Lambda-PFLOTRAN workflow for bulk OM simulations.  
378 To further assess and isolate the effect of OM speciation, extended forward simulations were performed by only  
379 varying FTICR-MS input data (Figure 7). FTICR-MS samples from Test Cases 1a-c and Test Case 2 were tested.  
380 These simulations replicate Figure 3 (i.e., Test Case 1a conditions and fitted  $\mu_{\text{max}}$  values) with the expectation of OM  
381 speciation, and demonstrate the significant impact of OM chemistry and speciation on overall predicted behavior,  
382 especially over longer time periods.  
383



384  
385 **Figure 7.** Influence of OM Speciation on Oxygen Consumption. FTICR-MS data from Test Cases 1a-c (grey shaded area), and  
386 Test Case 2 (blue line) were used as inputs. Bulk CH<sub>2</sub>O OM (green line) was also plotted for reference. Best fit  $\mu_{\text{max}}$  values to Test  
387 Case 1a were used (i.e., lambda binned  $\mu_{\text{max}} = 0.25 \text{ min}^{-1}$ ; bulk OM  $\mu_{\text{max}} = 0.032 \text{ min}^{-1}$ ).  
388

389 The clear variability in OM speciation, differences between a generic OM reaction network and one informed by  
390 FTICR-MS, and the impact of OM chemistry on biogeochemical predictive simulations underscore the importance of  
391 incorporating site-specific OM chemistry informed by ultra high resolution characterization into biogeochemical  
392 models.

## 393 5 Conclusions

394 Overall, Lambda-PFLOTRAN workflow provides an important linkage between molecular scale organic matter  
395 characterization and reactive transport simulations. This workflow allows for the influence of organic matter  
396 composition to be utilized within simulators to provide a more comprehensive understanding of the system chemistry

397 and behavior, moving beyond the standard assumption of bulk organic matter chemistry and composition. While there  
398 are current limitations due to how composition is characterized and quantified, this workflow connecting  
399 characterization information to simulations is an important advancement that can be refined as these laboratory  
400 techniques improve over time.

401 One of the major limitations surrounding this method, is the lack of understanding of organic matter compound  
402 bioavailability, resulting in a large conceptual gap as to how various organic carbon compounds may be utilized by  
403 microbes. In the absence of such information, all identified organic matter molecules are assumed to have equal  
404 bioavailability within this modeling framework when, in reality, compounds will exhibit varying degrees of  
405 bioavailability depending on factors such as associated size fraction, carbon pool, and environmental factors (Schmidt  
406 et al., 2011; Ahamed et al., 2023). Until improved understanding is established to discern individual compound  
407 bioavailability, this will remain as a limitation.

408 Another limitation of this method resides around the analytical limitations of organic carbon characterization and  
409 quantification. For instance, FTICR-MS focuses on water soluble organic matter which may provide a bias in the  
410 types of carbon identified by this technique (Tfaily et al., 2017). Additionally, as mentioned previously, FTICR-MS  
411 is qualitative, it does not provide structural information and will not differentiate between different isomers that have  
412 the same molecular formulas, it is only able to identify molecular formula is present or absent and not the concentration  
413 associated with each peak. Here, this has been addressed by assuming equal distribution of total carbon between the  
414 formulas within each  $\lambda$  bin on a per-carbon basis. This caveat can be easily updated in the workflow if new analytical  
415 advances are made that provide more quantitative information. Some existing approaches could be suitable for this  
416 type of modeling such as using quantitative biomarkers that cover major compound classes (Kim and Blair, 2023);  
417 but further advances in obtaining both high resolution and quantitative OM characterization would greatly aid in how  
418 we understand and model ecosystems.

419

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432  
433 **Code Availability:**

434 The source code, installation requirements, example test case notebooks, and associated data are available in ESS  
435 DIVE at <https://doi.org/10.15485/2281403>

436  
437 **Author Contribution:**

438 KM: conceptualization, formal analysis, methodology, software, writing- original draft preparation; PJ: methodology,  
439 software, writing- original draft preparation; GH: methodology, software, writing-review & editing; TA: data curation,  
440 software, writing-review & editing; HS: methodology, writing-review & editing; RK: supervision; NW: supervision,  
441 writing-review & editing; MB: investigation; RC: investigation; QZ: investigation; VG: investigation, data curation;  
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