Lambda-PFLOTRAN 1.0: Workflow for Incorporating Organic Matter Chemistry Informed by Ultra High Resolution Mass Spectrometry into Biogeochemical Modeling

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Abstract. Organic matter (OM) composition plays a central role in microbial respiration of dissolved organic matter and subsequent biogeochemical reactions. Here, a direct connection of organic ~~carbon~~ matter chemistry and thermodynamics to reactive transport simulators has been achieved through the newly developed Lambda-PFLOTRAN workflow tool that succinctly incorporates carbon chemistry data generated from Fourier transform ion cyclotron resonance mass spectrometry (FTICR-MS) into reaction networks to simulate organic matter degradation and the resulting biogeochemistry. Lambda-PFLOTRAN is a python-based workflow, executed through a Jupyter Notebook interface, that digests raw FTICR-MS data, develops a representative reaction network based on substrate-explicit thermodynamic modeling (also termed lambda modeling due to its key thermodynamic parameter λ used therein), and completes a biogeochemical simulation with the open source, reactive flow and transport code PFLOTRAN. The workflow consists of the following five steps: configuration, thermodynamic (lambda) analysis, sensitivity analysis, parameter estimation, and simulation output and visualization. Two test cases are provided to demonstrate the functionality of the Lambda-PFLOTRAN workflow. The first test case uses laboratory incubation data of temporal oxygen depletion to fit lambda parameters (i.e., maximum utilization rate and microbial carrying capacity). A slightly more complex second test case fits multiple lambda formulation and soil organic matter release parameters to temporal greenhouse gas generation measured during a soil incubation. Overall, the Lambda-PFLOTRAN workflow facilitates upscaling by using molecular-scale characterization to inform biogeochemical processes occurring at larger scales.

1 Introduction

Microbial respiration of dissolved organic carbon (DOC) is a main driver of environmental biogeochemical processes. Mechanistic biogeochemical models often rely on lumping organic matter into a few distinct carbon pools (e.g., dissolved, sorbed, mineral associated or refractory, labile, etc.) (e.g., Fatichi, et al., 2019, Robertson et al., 2019, Wang et al., 2013) but do not fully consider the properties of the organic matter (OM) compounds individually. Pooled carbon approaches have benefits, such as assigning variable levels of bioavailability, however, this approach does not capture the complex temporal dynamics of respiration driven by OM composition, as aerobic respiration rates have been linked to organic carbon concentration, thermodynamics of the OM (Stegen et al., 2018, Garayburu-Caruso et al., 2020), as well as the diversity of OM compounds present (Lehmann et al. 2020, Stegen et al., 2022). Such findings highlight the importance of incorporating individual OM chemistry into biogeochemical modeling to capture, and ultimately predict, system behavior more accurately.

There are many advanced instrumentation techniques capable of detecting and identifying individual OM formulae that comprise a bulk OM sample (e.g., GC-MS, HPLC-MS, Fourier transform ion cyclotron resonance mass spectrometry [FTICR-MS], etc.). For instance, FTICR-MS is a powerful, high-resolution, method that identifies molecular formulae for individual organic compounds. In any given environmental sample, FTICR-MS (or other ultra high-resolution methods) will typically resolve thousands of discrete OM molecular formulae, each with a unique mass and elemental composition (Cooper et al., 2020, Bahureksa et al., 2021). ~~Unfortunately~~However, untargeted analytical techniques like FTICR-MS are only able to determine if a compound is present and cannot quantify the total concentration associated with each organic matter molecule. Still, such techniques do provide immense amounts of characterization data encompassing a deeper analytical window than measuring a small number of individual biomarkers quantitatively (e.g., Ward et al., 2013). ~~However, the ability to~~ ~~U~~utilizinge such high-resolution molecular data in reactive transport modeling frameworks affords new opportunity to advance carbon cycling in terrestrial, riverine and coastal systems despite of ~~has remained a challenge and is typically not considered~~various theoretical and computational challenges.

Substrate-explicit thermodynamic modeling (SXTM) provides an avenue for incorporating individual OM reactivity based on thermodynamics (Song et al., 2020) into reactive transport models. The SXTM procedure takes the individual chemical formula derived from FTICR-MS (or another high-resolution technique) and uses its thermodynamic properties to generate an oxidation reaction for each molecular formula present in a sample. The corresponding reaction stoichiometry is then determined by considering catabolic, anabolic, and metabolic reactions and balancing energy for the overall metabolic reaction, allowing for the development of an aerobic respiration expression for each OM formula.

Still, the sheer number of compounds identified in each sample proves difficult for model integration. Typically, reactive transport simulators consider only a small number of primary species in their reaction networks, and most could not support modeling each of the thousands of organic matter molecules individually. Here, the developed Lambda-PFLOTRAN workflow addresses this challenge through grouping, or binning, similar compounds based on

their thermodynamic properties, allowing for the number of species considered within the reaction network to be reduced, and thus decreasing the required computational resources.

Lambda-PFLOTRAN is a python-based workflow that digests raw FTICR-MS data, develops a representative reaction network based on substrate-explicit thermodynamic modeling (Song et al., 2020), and completes a biogeochemical simulation with the open source, parallel reactive flow and transport code, PFLOTRAN (Hammond et al., 2014). PFLOTRAN is developed under an open source, GNU LGPL license. The term 'lambda' is used here because λ is a key parameter in the SXTM, which quantifies thermodynamic favorability of aerobic respiration of OM. The connection between the unique reaction network developed for each FTICR-MS sample hinges on the use of PFLOTRAN's reaction sandbox capability (Hammond, 2022). The reaction sandbox gives the ability to define additional custom, kinetic reactions beyond standard formulations (e.g., mineral precipitation-dissolution, Michaelis-Menten, etc.). The Lambda-PFLOTRAN workflow enables upscaling by using molecular-scale information to inform larger scale biogeochemical processes occurring throughout a watershed which can be simulated with PFLOTRAN. Herein we describe the Lambda-PFLOTRAN workflow process including the governing expressions, workflow steps, data requirements, as well as the associated assumptions and limitations. Two illustrative test cases are also included to demonstrate the ~~use of the~~ workflow~~. to utilize, parametrize, and model real datasets.~~

2 Methods

2.1 Conceptual Model

Respiration modeling herein is based on thermodynamic theory by Desmond-Le Quemener and Bouchez (2014) which was updated for multiple OM formulas by Song et al. (2020). The generalized form of OM molecule is assumed to take the form of $C_aH_bN_cO_dP_eS^z{}_f$. Each molecular formula then undergoes respiration (i.e., reaction with oxygen) based on the following general reaction expression:

$$y_{OM_i}OM_i + y_{H_2O}H_2O + y_{HCO_3^-}HCO_3^- + y_{NH_4^+}NH_4^+ + y_{HPO_4^{-2}}HPO_4^{-2-} + y_{HS^-}HS^- + y_{H^+}H^+ + y_{e^-}e^- + y_{O_2}O_2 + y_{BM}BM = 0, \quad (1)$$

This generalized expression is used to describe the oxidation of any OM molecule, *i*, and has been normalized to one mole of biomass (BM) produced. BM is assumed to have a formula of $CH_{1.8}O_{0.5}N_{0.2}$ (Stephanopoulos et al., 1998; Kleerebezem and Van Loosdrecht, 2010). OM_i represents the OM molecules as informed by FTICR-MS. Each *y* represents the reaction stoichiometry for that reactant ($y < 0$) or product ($y > 0$). While this expression is specific for cases where oxygen is the electron acceptor, such an expression could be updated for alternative electron acceptors.

Substrate-explicit thermodynamic modeling expressions developed from Song et al. (2020) were implemented in a reaction sandbox within PFLOTRAN. The expressions were implemented in a general manner allowing for flexibility in handling variations in FTICR-MS data and several user adjustable analysis configurations.

The microbial growth kinetics are described by Eq. (2):

$$\mu_i^{kin} = \mu^{max} exp(-\frac{\alpha|y_{OM\epsilon,i}|}{1000V_h[OM\epsilon,i]})exp(-\frac{\alpha|y_{O_2,i}|}{1000V_h[O_2]}), \quad (2)$$

where μ_i^{kin} is the unregulated uptake rate of reaction for OM_i [hr^{-1}], μ^{max} is the maximal microbial growth rate [hr^{-1}], $y_{OM,i}$ is the stoichiometry for OM_i [mol-OM · mol-biomass^{-1}], V_h is microbial harvest volume [m^3]. Given the physical interpretation of V_h as the microbial harvest volume, it is assumed here that the value of V_h is the same for both OM_i and O_2, [OM_i] is the organic matter concentration of OM_i [mol-OM·L^{-1}], $y_{O_2,i}$ is the stoichiometry for O_2 for respiration of OM_i [mol-O_2·mol-biomass^{-1}], [O_2] is oxygen concentration [mol-O_2·L^{-1}], α is a microbial unit conversion [mol-biomass] and 1000 is the conversion of m^3 to L.

Further, using a cybernetic modeling approach (after Song et al., 2018), all the unregulated uptake rates (μ_i^{kin}) are normalized by the sum of unregulated uptake rates across all reactions, *i* following Eq. (3):

$$u_i = \frac{\mu_i^{kin}}{\sum_{i=1}^{n} \mu_i^{kin}} \quad (3)$$

where u_i is the fraction of the unregulated rate [-]. The final regulated rate, r_i [hr^{-1}] for each reaction is then computed following Eq. (4):

$$r_i = u_i \mu_i^{kin}, \quad (4)$$

For implementation within PFLOTRAN, the use of inhibition terms was required to prevent negative concentrations once a reactant is nearly depleted. For a reaction to proceed, all reactant species must be present above a minimum concentration even if the molecules do not explicitly control the respiration rate (i.e., species other than OM and O_2, Eq. (2). If a reactant concentration falls below a threshold concentration, the respiration rate is inhibited. Reactant inhibition is computed by Eq. 5 (Kinzelbach et al., 1991) for reactant species *j:*

$$\boldsymbol{I_j = 0.5 + \frac{arctan([C_j]-C_{th_j})\cdot f}{\pi}}, \quad (5)$$

where $C_{th,i}$ is the threshold concentration [M], f is the threshold scaling factor [-]. The default C_{th_j} is 10^{-20} M.

The reaction rates are also inhibited by the microbial carrying capacity of the system, I_{cc}, as follows in Eq. (6):

$$I_{CC} = 1 - \frac{[BM]}{CC} \quad (6)$$

where [BM] is the biomass concentration [mol-BM·L^{-1}], CC is the biomass carrying capacity [mol-BM·L^{-1}]. I_{cc} has a non-negativity constraint, so if [BM] > CC, then I_{cc} = 0.

These inhibition factors are applied to the overall rate expression as shown in Eq. (7).

$$r_{i,inhibited} = r_i I_{CC} \prod I_j \quad \forall\, y_{i,j} < 0, \tag{7}$$

The overall individual species rates, d[C_j]/dt, [mol-species·L^{-1}·hr^{-1}] are then computed as follows with Eq. (8):

$$\frac{dC_j}{dt} = \left(\sum_{i=1}^{n} y_{i,j} r_{i,inhibited}\right)[BM], \tag{8}$$

where *j* is the species index. The total number of species includes 7 general species (i.e., HCO_3^-, NH_4^+, HPO_4^-, HS^-, H^+, O_2, BM (i.e., Eq (1)) and the OM species considered (i.e., typically 10). *i* is the reaction index, *n* is total number of reactions as based on the total number of OM species (typically, with this workflow *n* =10). $y_{i,j}$ is the coefficient for species *j* in reaction *i*.

The expression for biomass is also modified to account for biomass decay (note all biomass stoichiometries are 1 by definition):

$$\frac{dBM}{dt} = \left(\sum_{i=1}^{n} y_{i,j} r_{i,inhibited}\right)[BM] - k_{deg}[BM], \tag{9}$$

where k_{deg} is the biomass decay rate [hr^{-1}].

2.2 Lambda Analysis and Binning

To reduce the number of organic compounds considered in the simulation, OM molecules are grouped, or binned, based on their λ value computed by Eq. (10):

$$\lambda = \frac{\Delta G_{r,anabolic} + \Delta G_{r,dissipation}}{(-\Delta G_{r,catabolic})}, \tag{10}$$

where Δ*G* are the Gibbs energies for the anabolic and catabolic reactions and the associated dissipation energy, respectively. The value of λ is indicative of how many times the catabolic reaction needs to be completed to provide the energy required to synthesis one mole of biomass. Lower λ values suggest higher thermodynamic favorability of OM respiration. Using the chemical formula determined for each OM molecule, the energy balance equations are solved providing the overall reaction stoichiometry Eq. (1) and the λ is calculated. Using the λ value for each molecule, the cumulative probability distribution for the sample is produced (Figure 2).

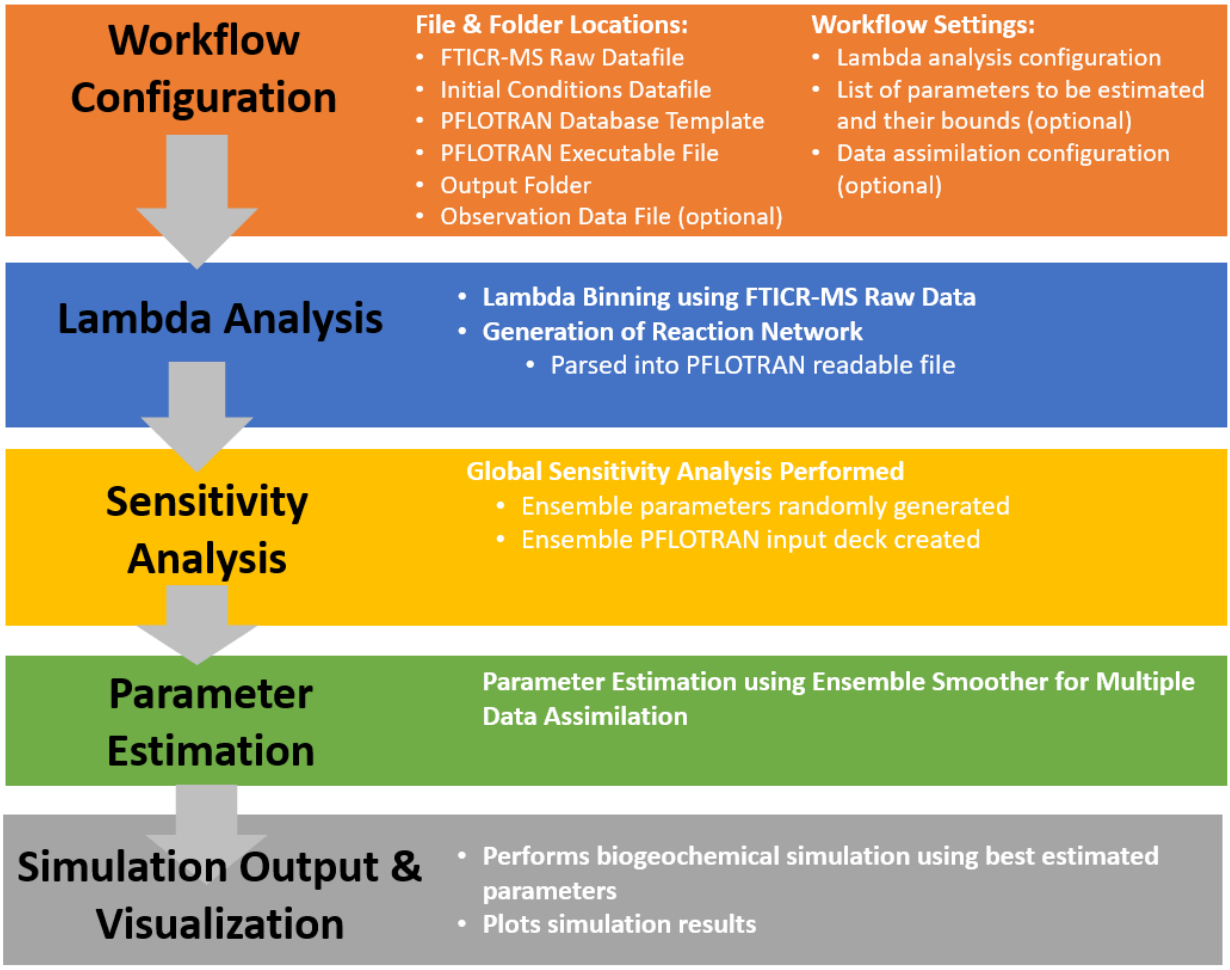


Figure 1: Flow Chart of the Lambda-PFLOTRAN Workflow.

It is this conversion from individual compounds to a distribution that is critical for reducing the entire sample down to a representative set of expressions. The λ bins are then formed by splitting the cumulative probability distribution into equally weighted sections as which to define the overall sample by. The illustrative example shown in Fig. 2 demonstrates the sample distribution being divided into 10 sections (i.e., in this case each section contains 10% of the overall sample distribution).

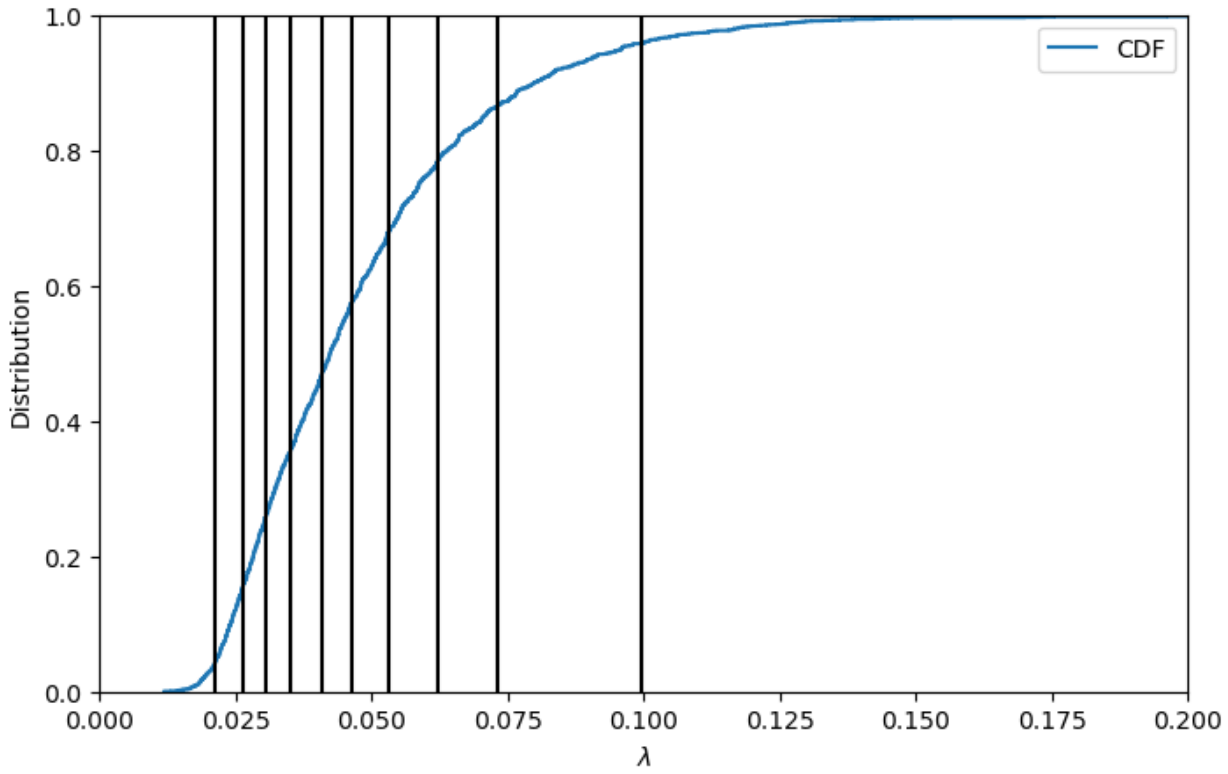


Figure 2: Lambda binning to convert raw FTICR-MS into a representative reaction network using the cumulative probability distribution function (CDF) for Test Case 1a. Vertical lines display the average λ value for each of the 10 bins (left to right, λ bin 1 to 10).

Each section is used to determine a representative organic matter formula and the associated reaction and stoichiometry of that λ bin. The group of representative reactions (one per bin) is called the reaction network. A demonstrative reaction network defined by λ analysis and binning is shown in Table 1.

Table 1: Reaction Network Developed from Lambda Theory for Test Case 1a

Bin Number	Representative Organic Matter Species Formula	λ	y_{OMC}	y_{HCO3^-}	y_{NH4^+}	$y_{HPO4^{-2}}$	y_{HS^-}	y_{H^+}	y_{O2}
1	$C_{31}H_{44}N_{0.33}O_{4.8}P_{0.6}S_{0.3}$	0.021	-0.05	0.64	-0.17	-0.18	0.03	0.02	-1.07
2	$C_{26}H_{39}N_{0.20}O_{7.0}P_{0.6}S_{0.1}$	0.026	-0.07	0.68	-0.10	-0.19	0.04	0.01	-1.06
3	$C_{22}H_{36}N_{0.24}O_{7.5}P_{0.5}S_{0.1}$	0.031	-0.08	0.69	-0.02	-0.18	0.04	0.01	-1.06
4	$C_{20}H_{32}N_{0.28}O_{7.3}P_{0.4}S_{0.1}$	0.035	-0.08	0.72	-0.08	-0.18	0.04	0.01	-1.05
5	$C_{19}H_{29}N_{0.48}O_{7.9}P_{0.3}S_{0.2}$	0.041	-0.09	0.79	-0.17	-0.16	0.03	0.02	-1.04
6	$C_{18}H_{26}N_{0.68}O_{8.1}P_{0.2}S_{0.2}$	0.046	-0.10	0.85	-0.27	-0.13	0.02	0.02	-1.03
7	$C_{17}H_{24}N_{0.69}O_{8.1}P_{0.2}S_{0.2}$	0.053	-0.11	0.90	-0.32	-0.12	0.02	0.02	-1.02
8	$C_{15}H_{20}N_{0.67}O_{7.6}P_{0.2}S_{0.2}$	0.062	-0.13	0.94	-0.42	-0.11	0.02	0.03	-1.00

9	$C_{13}H_{19}N_{1.13}O_{87.4}P_{0.1}S_{0.2}$	0.073	-0.15	1.01	-0.48	-0.03	0.01	0.03	-1.00
10	$C_{10}H_{15}N_{1.56}O_{6.5}P_{0.1}S_{0.2}$	0.100	-0.21	1.17	-0.75	0.12	0.01	0.04	-0.97

Currently, the representative OM molecule that defines each bin is computed as the average chemical formula of all the molecules present in that λ section. The disadvantage of this approach is that unrealistic compounds are defined as representative molecules instead of realistic molecules. The issue with selecting a single, but real compound, from within each λ section resides in chemical complexity and variation - for instance some molecules may contain low levels of phosphorous or sulfur and others may not contain either element in the chemical formula. Thus, requiring the representative chemical formula to be a real compound present in the sample would create basis which would propagate through the reaction network and into the resulting biogeochemical simulation results.

2.3 Lambda-PFLOTRAN Workflow

The Lambda-PFLOTRAN workflow digests raw FTICR-MS data, calculates the λ distribution for the sample, generates the λ bins and corresponding reaction network, and completes a biogeochemical simulation using PFLOTRAN. Further, we incorporated sensitivity analysis and ensemble data assimilation to enable an in-depth exploration of the impact of reaction parameters on respiration as well as a straightforward parameter estimation method to fit model parameters to experimental data.

The workflow is implemented through a user-friendly Jupyter notebook interface (Kluyver et al., 2016) where a user can configure the simulation parameters by adjusting initial concentrations, λ binning configuration, parameter values and/or ranges, and data assimilation options. Based on the user's data file and the associated parameters, scripts within the Jupyter notebook write the corresponding PFLOTRAN input files, including OM molecules and aqueous chemistry. The PFLOTRAN simulations are completed locally through a Docker container making this capability much more user-friendly and accessible. The progress of the data assimilation tool used for parameter fitting is illustrated within the Jupyter notebook. The resulting best fit final biogeochemical simulation is output visually with plots and as a text file (when applicable).

The Lambda-PFLOTRAN workflow steps are shown in Figure 1 and described in detail in the following subsections:

2.3.1 Step 1 – Workflow configuration

The first step is to set up the workflow configuration for a Lambda-PFLTORAN application. This includes specifying the file and folder locations of the following information: 1) FTICR-MS raw data file (.csv), 2) initial species concentrations file (.csv) that includes starting molar concentrations for HCO_3^-, NH_4^+, HPO_4^{2-}, HS^-, H^+, O_2 (aq), BM and total organic carbon (TOC), 3) PFLOTRAN database template file, 4) PFLOTRAN executable file, 5) workflow output folder, and if completing parameter estimation, (6) the data observation file (.csv), if applicable.

The user is also asked to configure workflow settings related to: (1) the lambda analysis configuration, including number of λ bins and method to define the λ bins (i.e., cumulative vs uniform); (2) the respiration modeling parameter

setup, including the list of the parameters to be estimated and their associated upper and lower bounds and (3) the data assimilation configuration (see below).

2.3.2 Step 2 – Organic Matter Chemistry using Lambda Analysis

With only an input of FTICR-MS data, the workflow first performs the lambda analysis (Section 2.2) to group OM molecules into various λ bins based on each compound's thermodynamics (Figure 2) and produce the corresponding reaction network for respiration (Table 1). The default number of λ bins is 10, although this can be adjusted in the workflow configuration by the user, if desired. The generated reaction network is then automatically parsed by the workflow into a text file that can be read by PFLOTRAN.

2.3.3 Step 3 – Sensitivity Analysis using Mutual Information

This step performs the global sensitivity analysis on the parameters to be estimated. Ensemble parameters are first generated by randomly sampling from their predefined ranges in the configuration step and saved into an HDF5 file. Then, the workflow generates a PFLOTRAN input deck to conduct ensemble simulations using the ensemble parameters. The generated ensemble model states enables a global sensitivity analysis using mutual information (Cover and Thomas, 2006; Jiang et al, 2022) as follows:

$$I(X;Y) = H(Y) - H(Y|X) = \sum_{X=x}\sum_{Y=y} p(x,y) log\left(\frac{p(x,y)}{p(x)p(y)}\right), \quad (11)$$

where x and y are the specific values of X and Y, respectively; $H(Y)$ is the Shannon's entropy of Y; $H(Y|X)$ is the conditional entropy of Y given X; p is the probability density function. Higher I indicate stronger sensitivity between X and Y. Besides sensitivity analysis, the ensemble parameter/states also serve as the prior information for parameter estimation at the next step.

2.3.4 Step 4 – Parameter Estimation using Ensemble Smoother for Multiple Data Assimilation

The workflow adopts Ensemble Smoother for Multiple Data Assimilation (Emerick and Reynolds, 2013; Jiang et al, 2021), abbreviated as ESMDA, for data assimilation in this step. Rooted in ensemble Kalman filter, ESMDA is an iterative data assimilation approach that assimilates the observations on the entire time period for multiple times to reduce the uncertainty of the estimated or posterior parameters. During each iteration of ESMDA, the model parameters are updated based on the following equation:

$$m_{k,l}^{u} = m_{k,l}^{f} + C_{MD,l}^{f}\left(C_{DD,l}^{f} + \alpha_l C_D\right)^{-1}\left(d_{obs} + \sqrt{\alpha_l} C_D^{\frac{1}{2}} z_k - d_{k,l}^{f}\right),\ k = 1, \dots, N_e \ and \ l = 1, \dots, L, \quad (12)$$

where the subscripts k and l are the indices of the ensemble member and the iteration, respectively; the superscripts u and f are the updated and forecast parameters or states, respectively; N_e is the number of ensemble members; L is the number of iterations; $m^f_{k,l}$ and $m^u_{k,l}$ are the kth ensemble member of the forecast/prior and updated/posterior parameters, respectively, at the lth iteration; d_{obs} is the observation; z_k is the observation noise sampled from

independent standard normal distributions for the kth ensemble member; $d^f_{k,l}$ is the kth ensemble member of the predicted observation states by the model using $m^f_{k,l}$; $C^f_{MD,l}$ is the cross-covariance matrix between the prior parameters m^f_l and the predicted observation states d^f_l; $C^f_{DD,l}$ is the auto-covariance matrix of the predicted observation states d^f_l; C_D is the auto-covariance matrix of the observation error; and α_l is the inflation coefficient at the lth iteration with the sum of all α_l equal to one.

Here, the assimilation starts with taking the ensemble model parameters/states in Step 3 and the provided observations, and calculates the posterior parameters using ensemble Kalman filter, updates the prior parameters with the current posterior for the next iteration, and then repeats the whole process for multiple times (typically 3 to 5 iterations, as defined by the user). The final estimated parameters are obtained from the posterior parameter at the last iteration and are updated in the parameter HDF5 file. The parameter estimation is implemented in a way that allows assimilating either a single (e.g., Test Case 1) or multiple observed species simultaneously through a simple change of the inputs. For example, if temporal experimental or field data is available for oxygen, pH, and total carbon, all these data sources could be simultaneously fit to, with only minor adjustments to Jupyter notebook.

2.3.5 Step 5 – Simulation Output and Visualization

The last step performs the ensemble simulation of the biogeochemical modeling a final time using the estimated parameters in Step 4. Optionally, users can further pick the realization with the best performance. The user has the option to select their preferred goodness of fit metric from the following options as a means for selecting the best performing simulation: R-squared (R^2), Root Mean Squared Error (RMSE), Modified Kling-Gupta Efficiency (m~~M~~KGE), Nash-Sutcliffe Model Efficiently Coefficient (NSE), or Correlation Coefficient (CorC). Based on the selection, the final time series of aqueous chemistry, oxygen consumption, CO_2 production, and lambda binned, and total organic carbon concentrations will be computed and plotted.

3 Test Cases

3.1 Test Case 1 - Oxygen Depletion Incubation Experiments.

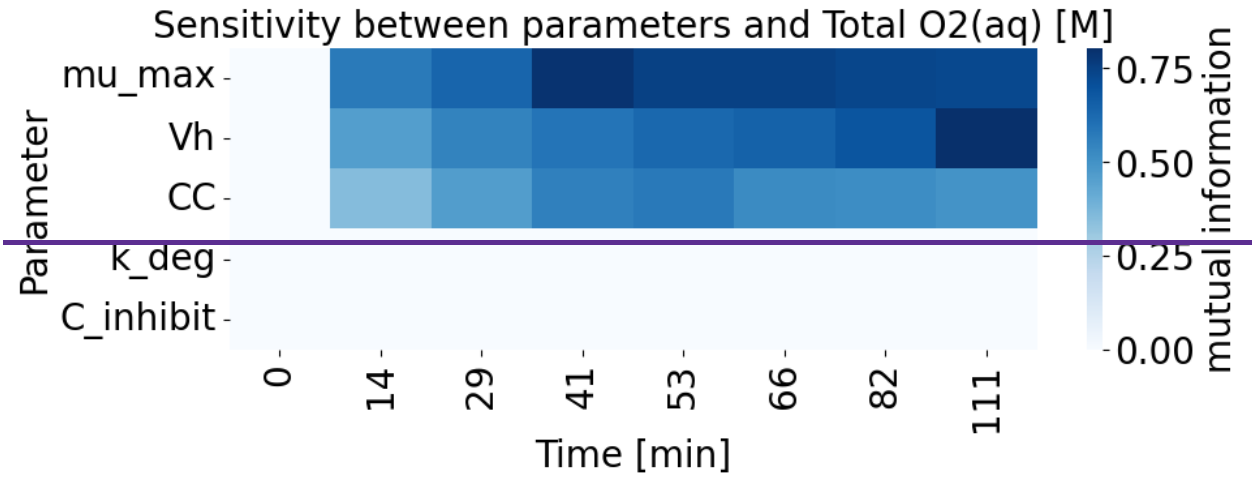
In the first illustrative example, the ~~lambda-pipelin~~workflow~~e~~ was used to fit ~~three-lambda-model-parameters (~~μ_{max}~~, V_h~~, and CC~~) to laboratory incubation experiments where oxygen levels were measured over two hours in a closed reactor. The incubation experiments were completed as part of the Worldwide Hydrobiogeochemistry Observation Network for Dynamic River Systems (WHONDRS) program (Goldman et al, 2020). For these incubations, sediment was taken from three locations within a stream (i.e., upstream [Test Case 1a], midstream [Test Case 1b], and downstream [Test Case 1c]) in the Yakima River Basin in Washington, USA for subsequent laboratory respiration experiments. FTICR-MS was used to determine the OM chemistry from each sediment sample, resulting in variable formulae being identified in each sample. Formula assignments for all the samples included herein were completed using formultitude~~formularity~~ (Tolic et al., 2017). Total dissolved organic carbon concentration paired with the FTICR-MS sample and biomass measurements taken at the start of each experiment were used as the initial

concentrations for each of the simulations. Due to the absence of quantitative data related to how the total carbon mass is distributed between various the OM compounds, the total carbon concentration (on a per-C basis) was assumed to be split equally between each of the λ bins. The total organic carbon concentration was distributed into each λ bin using Eq. (13). While this assumption results in equal distribution of carbon between the bins, consequently, it assigns different initial species concentrations due to varying carbon concentrations between the molecules.

$$[C_{\lambda \mathrm{bin}}]_0 = \frac{[TOC]}{n_{\lambda bin} nC_{\lambda bin}} \quad (13)$$

Where: $[C_{\lambda \mathrm{bin}}]_0$ is the initial species concentration in each λ bin [mol·L^{-1}]; TOC is the total organic carbon measured [mol-carbon·L^{-1}]; $n_{\lambda bin}$ is the number of λ bins [-]; and $nC_{\lambda bin}$ is the number of carbon molecules in the assumed formula for the λ bins [mol-carbon · mol-molecule^{-1}].

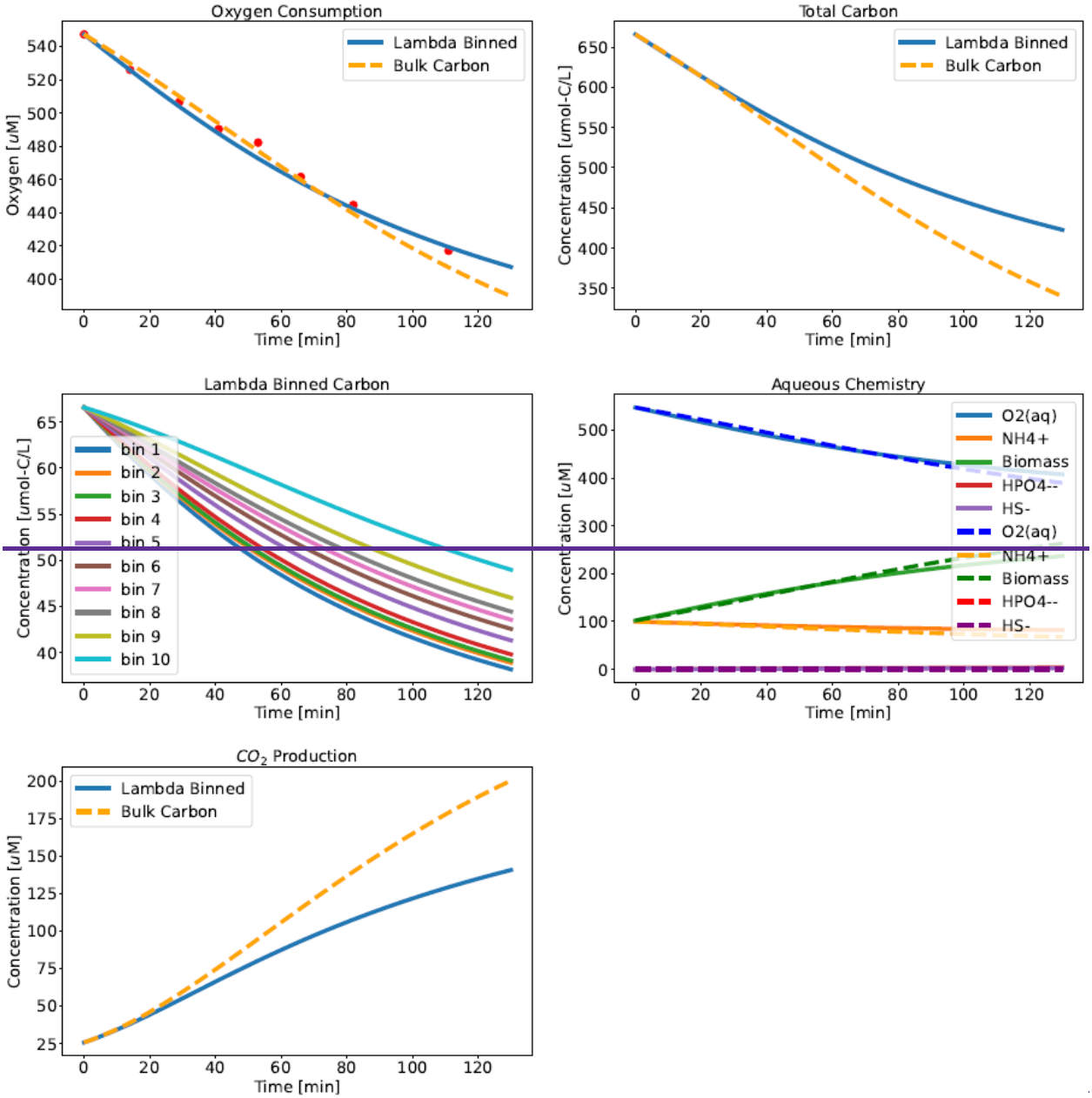
Using the Lambda-PFLOTRAN workflow, the FTICR-MS data from each laboratory experiment was digested into the corresponding λ bins to create the individual reaction network. The Jupyter Notebook for this example is "Test_Case1-WHONDRS.ipynb" and is available at https://doi.org/10.15485/2281403. ~~The resulting λ binning and associated reaction network for Test Case 1a are shown in Figure 2 and Table 1. Test cases 1b and 1c are in the Supporting Information (Fig. S1 – S2 and Tables S2 – S3). The calculated parameter sensitivity is shown in Figure 3, which indicates the results highly sensitive to all three parameters, in particular μ_{max} and V_h more so than and CC.~~



~~**Figure 3:** Illustrative Example of Sensitivity Analysis Output during Parameter Estimation. Example shown here provides the sensitivity of three fitted parameters (μ_{max}, V_h, and CC) on temporal aqueous O_2 concentrations as a function of time.~~

μ_{max} ~~Lambda expression parameters were~~ was fit to the provided experimental oxygen data. The final lambda binned fit, along with ~~and the final fit to the experimental data and~~ corresponding carbon consumption (individual and total) and aqueous chemistry is displayed in Figure ~~4~~ 3 (and in the supporting information Fig. ~~S2~~ S1 and ~~S3~~ S2 for Test

Cases 1b (midstream) and 1c (downstream), respectively). ~~The workflow was also run where μ_{max} was fit again, but this time assuming a generic OM form of CH_2O. This, allowing~~ To evaluate the ~~allows for comparison between using information for~~use of lambda binned OM obtained from FTICR-MS (Figure ~~4~~3), ~~,~~ the workflow was also run for a baseline case where μ_{max} was fit again, but this time assuming a generic bulk OM form of CH_2O for comparison. Fitted μ_{max} values for the lambda binned model is 0.25 min^{-1} (R^2 = 0.99) and fitted μ_{max} to the bulk OM CH_2O model is 0.032 min^{-1} (R^2 = 0.96). V_h and CC are fixed at assumed values of 10 m^3 and 1 M, respectively in both simulations. ~~In this case, the same set of lambda parameters were fit to the oxygen consumption experimental data, which also resulted in successful fit (R^2 = 0.990 for lambda binned model; R^2 = 0.987 for bulk model).~~



~~**Figure 4:** Test Case 1a Results – Oxygen Consumption (top left) where Lambda-PFLOTRAN workflow was used to fit (blue line) to experimental respiration data (red dots) and the corresponding Total Carbon Consumption (top right); Individual Organic Matter Consumption by λ bin (middle left); corresponding biogeochemistry including O_2 (aq) (blue); Biomass (green); NH_4^+ (orange); HS^- (purple); and HPO_4^{--} (red) (middle left); and CO_2 production (bottom left) for the upstream incubation. The dashed orange lines in the top two figures show simulation results assuming a generic OM species of CH_2O for comparison. Fitted values for the lambda binned model are μ_{max} = 0.25 min^{-1}, V_h = 9.7 m^3, and CC = 0.49 M (R^2 = 0.990), and fitted bulk OM CH_2O model value are μ_{max} = 0.05 min^{-1}, V_h = 3.3 m^3, and CC = 0.58 M (R^2 = 0.987).~~

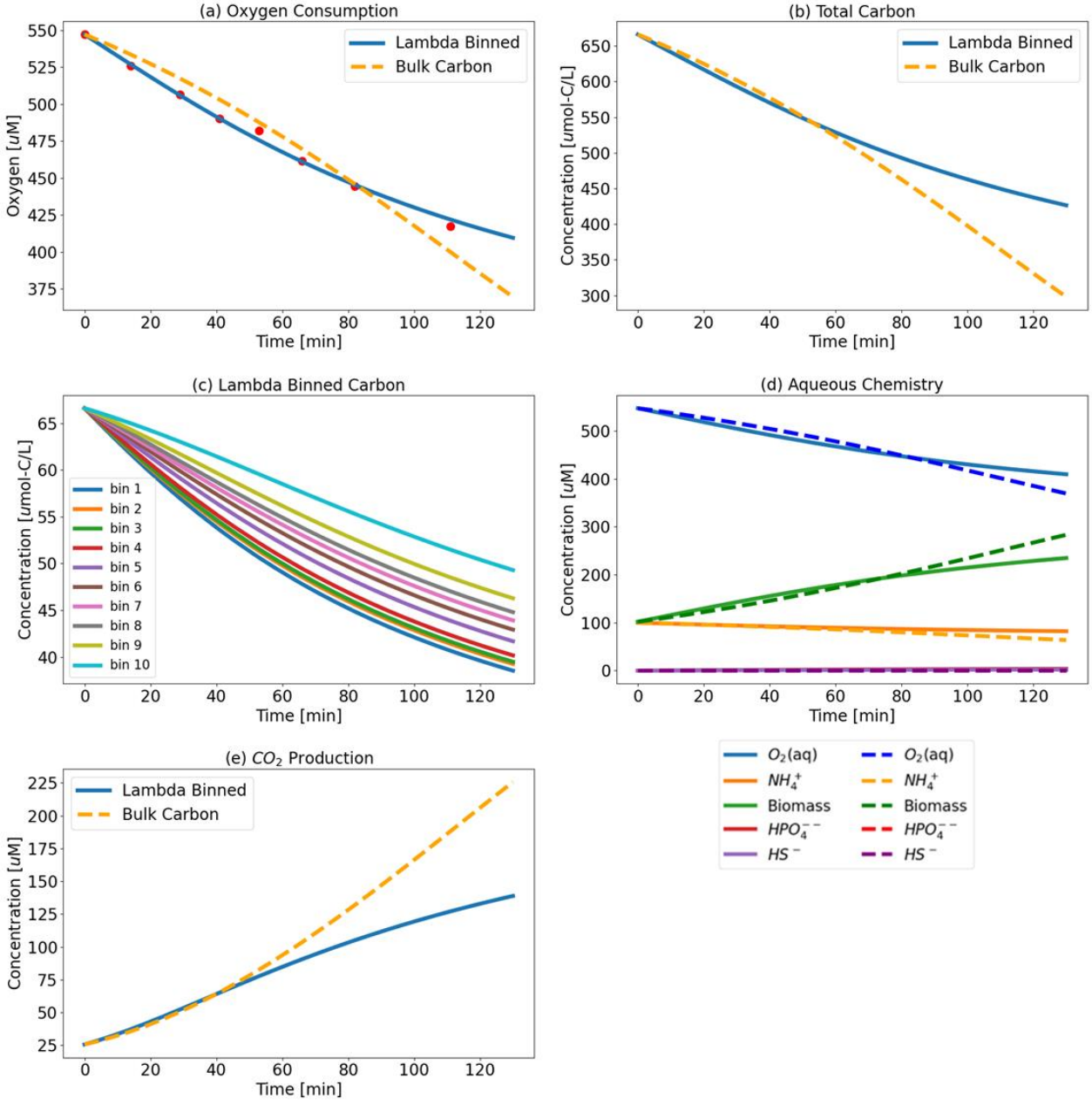


Figure 3: Test Case 1a Results – (a) Oxygen Consumption where Lambda-PFLOTRAN workflow was used to fit (blue line) to experimental respiration data (red dots) and (b) Total Carbon Consumption; (c) Individual Organic Matter Consumption by λ bin; and (d) biogeochemistry including O_2 (aq) (blue); Biomass (green); NH_4^+ (orange); HS^- (purple); and HPO_4^{--} (red); and (e) CO_2

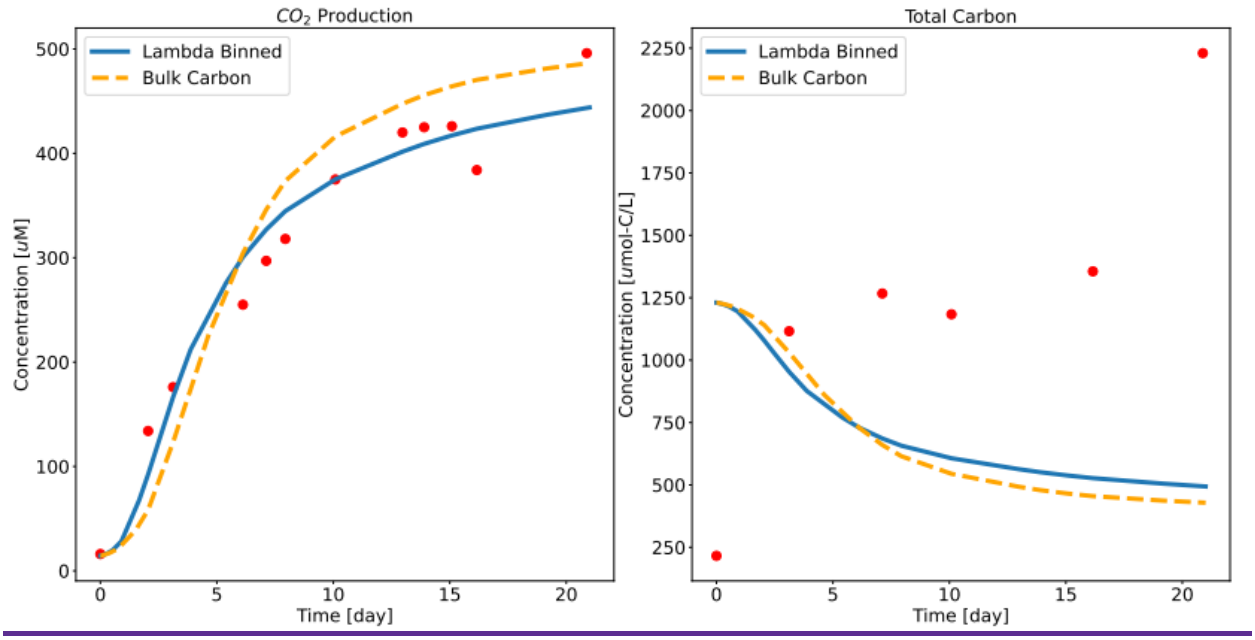
production for the upstream incubation. The dashed orange lines (in a, b and e) show simulation results assuming a generic OM species of CH_2O for comparison.

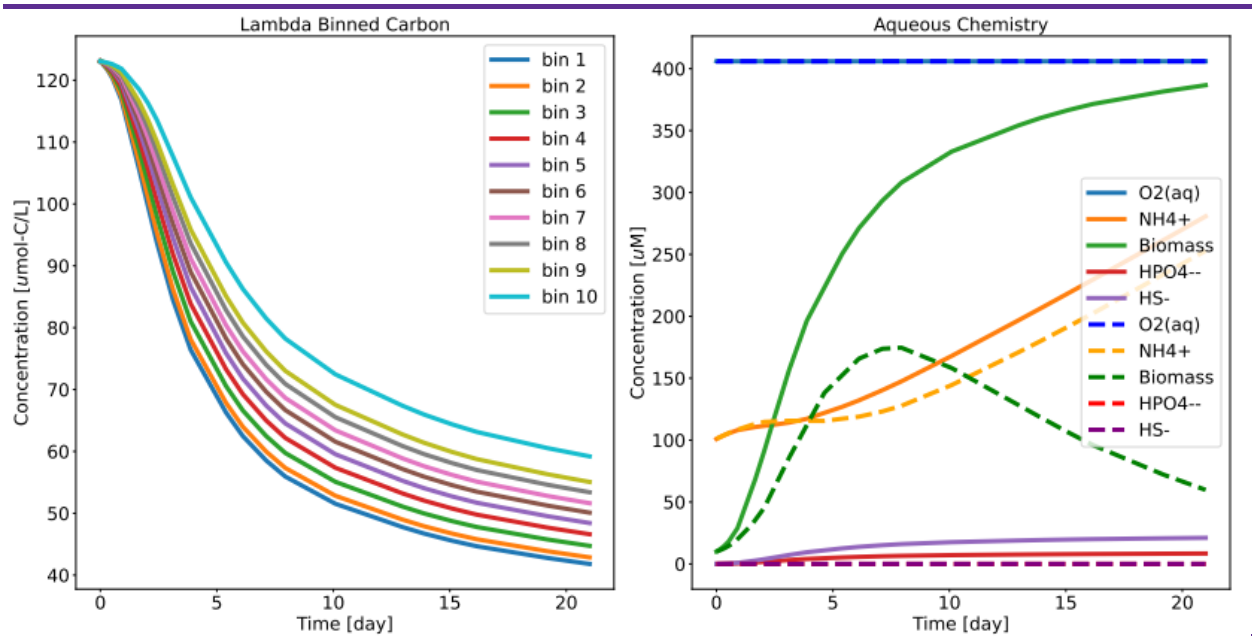
However, even over the short time frame of this simulation (i.e., only 120 minutes), the difference between assuming the generic CH_2O and using the more detailed organic matter chemistry resulted in different predictions of total carbon and CO_2 generation. The bulk OM model predicts more carbon consumption and greater CO_2 production than the ~~binned~~ lambda binned model. The bulk OM model estimates that ~~65~~50% of the initial total carbon is consumed over the first 120 mins, whereas the lambda binned model predicts ~~54~~34% consumption. Similarly, the bulk OM model estimates approximately ~~41~~35% more CO_2 generation as compared the lambda binned model. The effects on aqueous chemistry over this short duration are more muted, albeit still present.

3.2 Test Case 2 - Respiration Incubation Experiments.

Test Case 2 uses soil respiration incubation data from Ward et al. (2023) aimed at investigating the influence of soil type, oxygen condition (aerobic vs. anaerobic), and seawater exposure (fresh vs. saline) on respiration extent and rate. For these experiments, temporal measurements were collected for CO_2 generation, dissolved organic carbon (DOC), organic matter formulas via FTICR-MS and other bulk aqueous chemistry (i.e., pH, NH_4^+, and other metals and ions) creating a rich dataset for calibration of system specific lambda model parameters. These incubations were setup by adding dry soil to the reactor and then adding water (resulting in a soil:water ratio ranging from 1:11 to 1:16). The soil and water were shaken vigorously for five minutes, and then sampled for the initial time point prior to officially starting the incubation. For the aerobic experiments, the reactor headspace was cycled every 24 hours to measure CO_2 generated but also to ensure the system was kept aerobic; this was only performed five days per week, with no measurements taken on the weekend due to logistical constraints. Upon experiment completion, the increase in DOC concentrations indicated organic carbon was being kinetically released from the soil into the aqueous phase over the course of the 21-day experiment. Similarly, measured NH_4^+ concentrations also increased during the experiment. To address this within our reactive transport model, a source of nitrogen was assumed to be released from the soil as well ($N_{release}$). Both carbon and nitrogen release are included in this example and are assumed to follow a zero-order constant release rate. Any organic carbon released from the soil was fractionated into each λ bin on the same per-carbon basis assumed for the initial total organic carbon. This was implemented through a dependent function that calculated the release of carbon into each λ bin based on a fitted single bulk $k_{release}$ rate. Mathematically in PFLOTRAN the constant oxygen conditions were implemented through a gas-liquid partitioning expression with a fast exchange term. These three additional processes were added to describe the experimental conditions of Test Case 2 more accurately (i.e., release of carbon, nitrogen and sustained aerobic conditions); however, a PFLOTRAN input deck can be expanded and customized to include a host of additional processes and full geochemistry for a specific system of interest. For instance, aqueous complexation, mineral dissolution and precipitation, sorption, and redox reactions can be added, all of which can influence the resultant pH and carbon, nitrogen, and other nutrient dynamics.

The workflow was used to fit μ_{max}, V_h, CC, k_{deg}, as well as $k_{release}$, to the temporal CO_2 generation for a single aerobic soil incubation (Figure 5). The Jupyter Notebook for this example is “Test_Case2-Colloids.ipynb”.


CO_2 Production
Lambda Binned
Bulk Carbon
Concentration [uM]
Time [day]
Total Carbon
Lambda Binned
Bulk Carbon
Concentration [umol-C/L]
Time [day]


Lambda Binned Carbon
bin 1
bin 2
bin 3
bin 4
bin 5
bin 6
bin 7
bin 8
bin 9
bin 10
Concentration [umol-C/L]
Time [day]
Aqueous Chemistry
O2(aq)
NH4+
Biomass
HPO4--
HS-
O2(aq)
NH4+
Biomass
HPO4--
HS-
Concentration [uM]
Time [day]

336

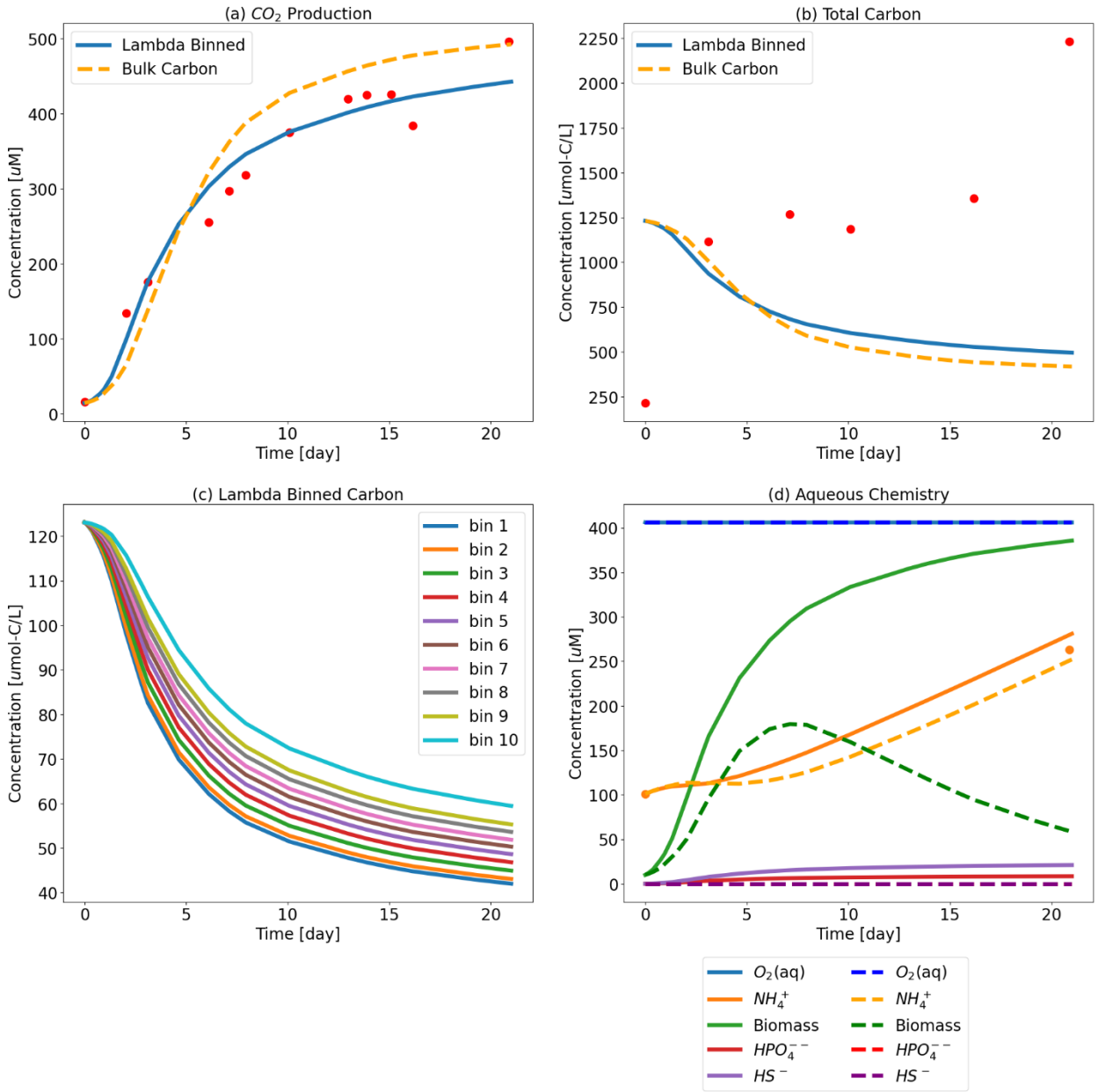


Figure ~~5~~4. Test Case 2 Results – (a) CO_2 production ~~top left~~ where Lambda-PFLOTRAN workflow was used to fit (blue line) to experimental respiration data (red dots) and (b) the corresponding Total Organic Carbon ~~top right~~; (c) Individual Organic Matter Consumption by λ bin ~~bottom left~~ and (d) the corresponding biogeochemistry including O_2 (aq) (blue); Biomass (green); NH_4^+ (orange); HS^- (purple); and HPO_4^{--} (red) ~~bottom right~~. Dots indicate experimental data. The dashed orange lines in the top two figures show simulation results assuming a generic OM species of CH_2O for comparison. Fitted parameters for lambda binned model were $k_{release}$ = 5.5x10^{-12} day^{-1}; μ_{max} = 37.6 day^{-1}, V_h = 5.0 m^3, CC = 0.12 M, and k_{deg} = 1x10^{-3} day^{-1} (R^2 = 0.953) and fitted bulk OM CH_2O model values were $k_{release}$ = 2.0x10^{-12} day^{-1}; μ_{max} = 47 day^{-1}, V_h = 1.0 m^3, CC = 0.77 M, and k_{deg} = 0.15 day^{-1} (R^2 = 0.909).

~~The best fit results indicate a superior fit using the lambda binned OM over the bulk OM model and in fact, the bulk model is unable to successfully capture the temporal evolution of the CO_2. It should be noted that both model fits are~~

~~highly sensitive to the allowable parameter space as user defined by the lower and upper parameter bounds.~~ For the purposes for showcasing the workflow, five parameters were estimated in this test case example, and as a result the models are ~~likely~~ over parametrized given the amount of data available. Parameter sensitivity over the course of simulation time is shown in Figure 5 and suggests this system is highly sensitive to V_h. It should be noted that both these model fits are also highly sensitive to the allowable parameter space as user defined by the lower and upper parameter bounds.

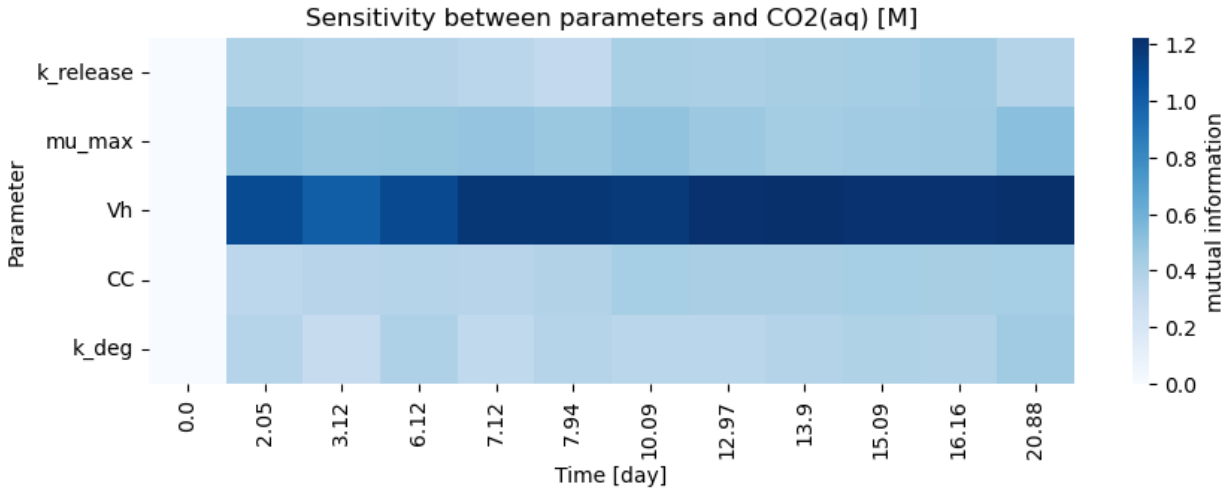


Figure 5. Test Case 2 - Sensitivity Analysis Output during Parameter Estimation. The sensitivity of five fitted parameters ($k_{release}$, μ_{max}, V_h, CC, and k_{deg}) on temporal aqueous CO_2 concentrations as a function of time.

Any additional experimental data, either collected during incubations or through independent experiments (e.g., carbon release from the soil in an abiotic system), would be expected to help constraint the model and improve parameterization. Additionally, it is unclear why the model is unable to capture the total organic carbon behavior in Test Case 2. One potential explanation is that some of the released organic carbon may not be fully bioavailable and thus the model may be compensating for this by artificially reducing the concentration of OM available for respiration.

4 Variability and Impact of Organic Matter Speciation

The variability in OM speciation was briefly assessed by comparing FTICR-MS data from Test Cases 1 and 2. Each identified OM species was classified into one of nine compound classes. For Test Case 1, the average of the three Test Case 1 samples (1a - upstream, 1b - midstream, and 1c - downstream) was computed. The predominant classes were proteins (34 ± 1%), lignin (26 ± 1%), and lipids (13 ± 2%), with the errors representing the standard deviation among the Test Case 1a-c samples. The low standard deviation suggests consistent reproducibility in OM speciation for samples taken from nearby locations. In contrast, OM in Test Case 2 was primarily composed of lignin (37.4%) and concentrated hydrocarbons (32%). The full distribution of compound classes is presented in Figure 5.

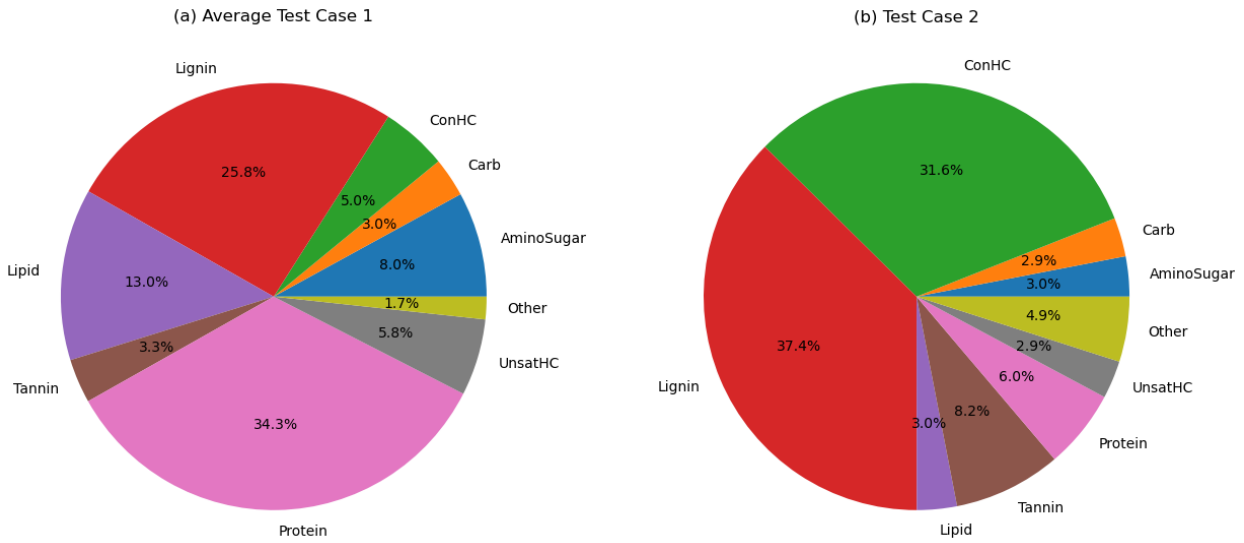


Figure 5. Distribution of Organic Matter Compound Classes: (a) Test Case 1 and (b) Test Case 2.

Note: Test Case 1 is the average of Test Case samples 1a-c. ConHC = Condensed Hydrocarbon; UnsatHC = Unsaturated Hydrocarbon

The influence of the sample OM speciation on the λ binned reaction networks was also assessed. Figure 6 illustrates the impact of OM speciation on the corresponding λ binned reaction networks, with three key observations. First, the variability in OM speciation between different samples is evident when comparing Test Case 1 and Test Case 2. To enhance visual clarity, the range of Test Case 1 samples (1a-c) is depicted as a grey shaded region, showing the spread between the minimum and maximum values of the three samples. For Test Case 2, data from the single FTICR-MS sample is represented by blue dots. Test Case 1 and 2 have distinct λ derived reaction networks as indicated by the little overlap between the grey region and the blue dots in Figures 6b-i.

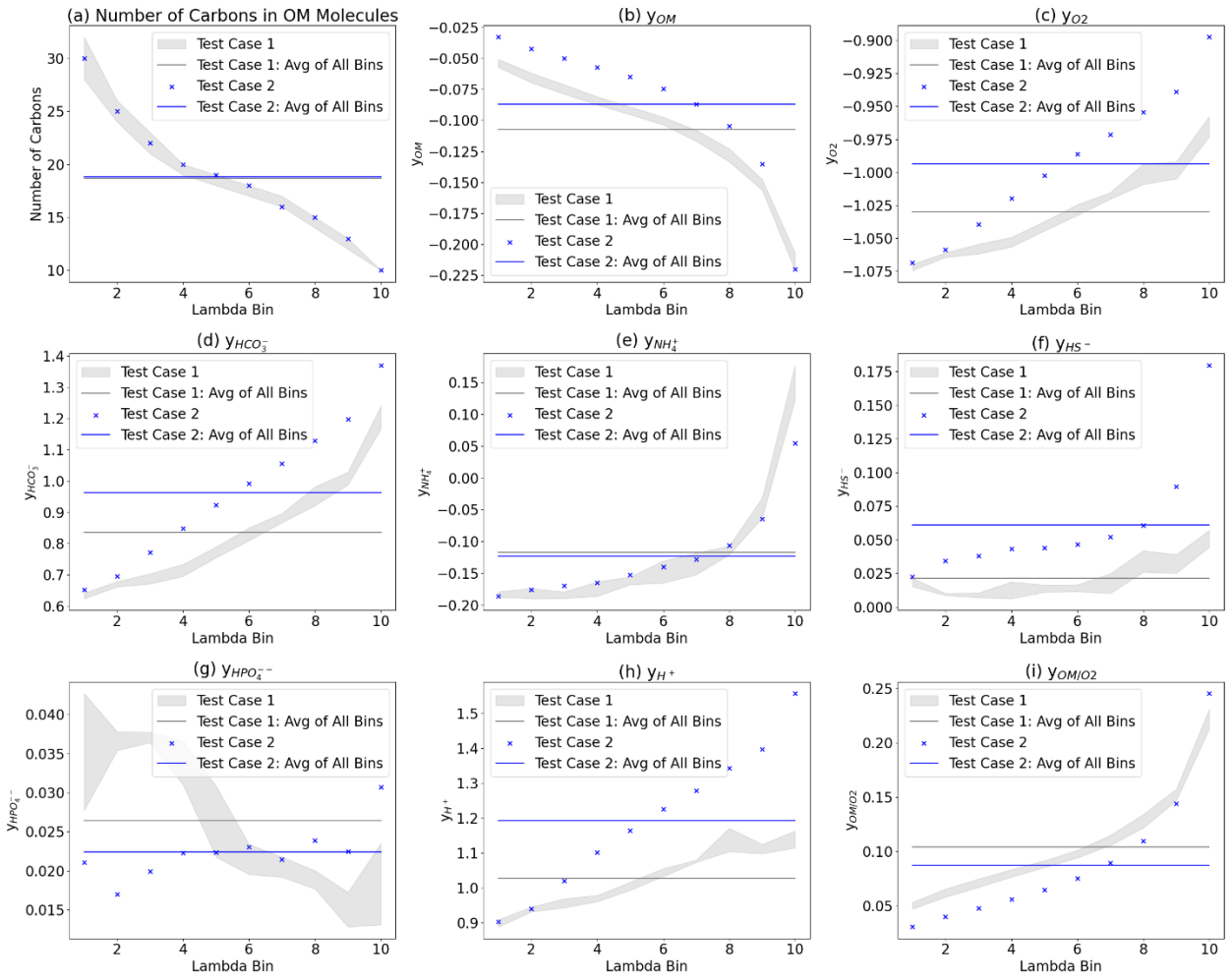


Figure 6. Comparison of Lambda-Binned Reaction Network Parameters: (a) number of carbons in the OM; stoichiometric coefficient, *y*, for (b) OM, (c) O_2, (d) HCO_3^-, (e) NH_4^+, (f) HS^-, (g) HPO_4^{--}, (h) H^+; and (i) ratio of OM/O_2 coefficients for Test Case 1a-c (grey dots); the average of all λ bins for Test Case 1 (grey line); Test Case 2 (blue x); and the average of all λ bins for Test Case 2 (blue line). The grey shaded area highlights the range of values for Test Case 1a-c for better visual comparison.

Second, the λ binning process captures the OM speciation variation within a sample. To illustrate this intrasample variability, a line representing the average of all λ bins is shown on Figure 6 (grey line for Test Case 1, blue line for Test Case 2). The difference between the reaction network coefficients (vertical axis) for the λ binning (grey shaded area and blue dots) and the Test Case average lines highlights the extent of this variability. Finally, although the λ binning process resulted in a similar number of carbon atoms to OM molecules within each λ bin for both test cases (Figure 6a), the resulting stoichiometric coefficients in the reaction networks differ significantly (Figures 6b-h). These stoichiometric differences lead to variations in biogeochemical outcomes, such as OM-to-oxygen utilization ratios during aerobic respiration (Figure 6i). These differences are due to the additional elements beyond carbon in the OM molecules (i.e., nitrogen, oxygen, sulfur, hydrogen, and phosphorus).

Another important aspect is the comparison against assuming a generic, bulk OM composition, which does not account for OM speciation as informed by FTICR-MS or similar methods. The reaction network developed for a generic OM molecule of CH_2O is shown in Eq. 14.

$$2.03\ CH_2O + 0.98\ O_2 + 0.2\ NH_4^+ \rightarrow 1.03\ HCO_3^- + 1.23\ H^+ + 0.4\ H_2O + CH_{1.8}O_{0.5}N_{0.2} \quad (14)$$

This reaction network is used in the Lambda-PFLOTRAN workflow for bulk OM simulations.

To further assess and isolate the effect of OM speciation, extended forward simulations were performed by only varying FTICR-MS input data (Figure 7). FTICR-MS samples from Test Cases 1a-c and Test Case 2 were tested. These simulations replicate Figure 3 (i.e., Test Case 1a conditions and fitted μ_{max} values) with the expectation of OM speciation, and demonstrate the significant impact of OM chemistry and speciation on overall predicted behavior, especially over longer time periods.

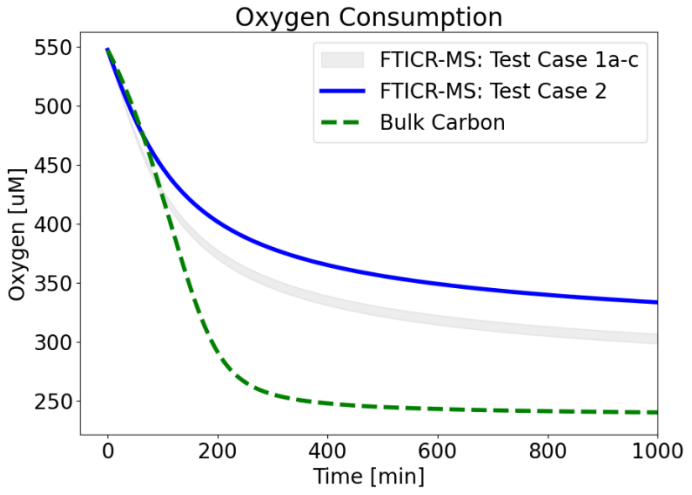


Figure 7. Influence of OM Speciation on Oxygen Consumption. FTICR-MS data from Test Cases 1a-c (grey shaded area), and Test Case 2 (blue line) were used as inputs. Bulk CH_2O OM (green line) was also plotted for reference. Best fit μ_{max} values to Test Case 1a were used (i.e., lambda binned μ_{max} = 0.25 min^{-1}; bulk OM μ_{max} = 0.032 min^{-1}).

The clear variability in OM speciation, differences between a generic OM reaction network and one informed by FTICR-MS, and the impact of OM chemistry on biogeochemical predictive simulations underscore the importance of incorporating site-specific OM chemistry informed by ultra high resolution characterization into biogeochemical models.

~~4~~ 5 Conclusions

Overall, Lambda-PFLOTRAN workflow provides an important linkage between molecular scale organic matter characterization and reactive transport simulations. This workflow allows for the influence of organic matter composition to be utilized within simulators to provide a more comprehensive understanding of the system chemistry

and behavior, moving beyond the standard assumption of bulk organic matter chemistry and composition. While there are current limitations due to how composition is characterized and quantified, this workflow connecting characterization information to simulations is an important advancement that can be refined as these laboratory techniques improve over time.

One of the major limitations surrounding this method, is the lack of understanding of organic matter compound bioavailability, resulting in a large conceptual gap as to how various organic carbon compounds may be utilized by microbes. In the absence of such information, all identified organic matter molecules are assumed to have equal bioavailability within this modeling framework when, in reality, compounds will exhibit varying degrees of bioavailability depending on factors such as associated size fraction, carbon pool, and environmental factors (Schmidt et al., 2011; Ahamed et al., 2023). Until improved understanding is established to discern individual compound bioavailability, this will remain as a limitation.

Another limitation of this method resides around the analytical limitations of organic carbon characterization and quantification. For instance, FTICR-MS focuses on water soluble organic matter which may provide a bias in the types of carbon identified by this technique (Tfaily et al., 2017). Additionally, as mentioned previously, FTICR-MS is qualitative, it does not provide structural information and will not differentiate between different isomers that have the same molecular formulas, it is only able to identify molecular formula is present or absent and not the concentration associated with each peak. Here, this has been addressed by assuming equal distribution of total carbon between the formulas within each λ bin on a per-carbon basis. This caveat can be easily updated in the workflow if new analytical advances are made that provide more quantitative information. Some existing approaches could be suitable for this type of modeling such as using quantitative biomarkers that cover major compound classes (Kim and Blair, 2023); but further advances in obtaining both high resolution and quantitative OM characterization would greatly aid in how we understand and model ecosystems.

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Code Availability:

The source code, installation requirements, example test case notebooks, and associated data are available in ESS DIVE at https://doi.org/10.15485/2281403

Author Contribution:

KM: conceptualization, formal analysis, methodology, software, writing- original draft preparation; PJ: methodology, software, writing- original draft preparation; GH: methodology, software, writing-review & editing; TA: data curation, software, writing-review & editing; HS: methodology, writing-review & editing; RK: supervision; NW: supervision, writing-review & editing; MB: investigation; RC: investigation; QZ: investigation; VG: investigation, data curation; AR: investigation; XC: conceptualization, investigation, writing-review & editing

Competing Interests: The authors declare that they have no conflict of interest.

References

Bahureksa, W., Tfaily, M. M., Boiteau, R. M., Young, R. B., Logan, M. N., McKenna, A. M., & Borch, T. (2021). Soil organic matter characterization by Fourier transform ion cyclotron resonance mass spectrometry (FTICR MS): A critical review of sample preparation, analysis, and data interpretation. *Environmental science & technology*, 55(14), 9637-9656, https://doi.org/10.1021/acs.est.1c01135

Cover, T. M., and Thomas, J. A. (2006). Elements of information theory (Wiley series in telecommunications and signal processing). Wiley-Interscience.

Emerick, A.A., Reynolds, A.C. (2013). Ensemble smoother with multiple data assimilation. *Comput. Geosci*. 55, 3–15. https://doi.org/10.1016/j.cageo.2012.03.011.

Fatichi, S., Manzoni, S., Or, D., & Paschalis, A. (2019). A mechanistic model of microbially mediated soil biogeochemical processes: a reality check. *Global Biogeochemical Cycles*, *33*(6), 620-648. https://doi.org/10.1029/2018GB006077

Garayburu-Caruso, V., Stegen, J., Song, H.-S., Renteria, L., Wells, J., Garcia, W., et al. (2020). Carbon limitation leads to thermodynamic regulation of aerobic metabolism. *Environ. Sci. Technol. Lett*. 7, 517–524. https://doi.org/10.1021/acs.estlett.0c00258

Goldman A E ; Arnon S ; Bar-Zeev E ; Chu R K ; Danczak R E ; Daly R A ; Delgado D ; Fansler S ; Forbes B ; Garayburu-Caruso V A ; Graham E B ; Laan M ; McCall M L ; McKever S ; Patel K F ; Ren H ; Renteria L ; Resch C T ; Rod K A ; Tfaily M ; Tolic N ; Torgeson J M ; Toyoda J G ; Wells J ; Wrighton K C ; Stegen J C ; WHONDRS Consortium T (2020): WHONDRS Summer 2019 Sampling Campaign: Global River Corridor Sediment FTICR-MS, Dissolved Organic Carbon, Aerobic Respiration, Elemental Composition, Grain Size, Total Nitrogen and Organic Carbon Content, Bacterial Abundance, and Stable Isotopes (v8). River Corridor and Watershed Biogeochemistry SFA, ESS-DIVE repository. Dataset. doi:10.15485/1729719 accessed via https://data.ess-dive.lbl.gov/datasets/doi:10.15485/1729719 on 2023-12-28

Hammond, G. E., Lichtner, P. C., & Mills, R. T. (2014). Evaluating the performance of parallel subsurface simulators: An illustrative example with PFLOTRAN. *Water resources research*, *50*(1), 208-228. https://doi.org/10.1002/2012WR013483

Hammond, G.E. (2022) The PFLOTRAN Reaction Sandbox, *Geoscientific Model Development*, 15, 1659-1676, https://doi.org/10.5194/gmd-15-1659-2022.

Jiang P., Chen, X., Chen, K., Anderson, J., Collins, N., Gharamti, M. (2021) DART-PFLOTRAN: An ensemble-based data assimilation system for estimating subsurface flow and transport model parameters. *Environmental Modelling & Software*, Volume 142, https://doi.org/10.1016/j.envsoft.2021.105074.

Jiang, P., Son, K., Mudunuru, M.K. and Chen, X. (2022). Using mutual information for global sensitivity analysis on watershed modeling. *Water Resources Research*, 58(10), https://doi.org/10.1029/2022WR032932

Kim, J., Blair, N.E. Biomarker heatmaps: visualization of complex biomarker data to detect storm-induced source changes in fluvial particulate organic carbon. Earth Sci Inform 16, 2915–2924 (2023). https://doi.org/10.1007/s12145-023-01039-y

Kinzelbach, W., Schafer, W., and Herzer, J. (1991). Numerical modeling of natural and enhanced denitrification processes in aquifers. *Water Resources Research*, 27(6):1123–1135. https://doi.org/10.1029/91WR00474

Kleerebezem, R., & Van Loosdrecht, M. C. (2010). A generalized method for thermodynamic state analysis of environmental systems. Critical Reviews in Environmental Science and Technology, 40(1), 1-54. https://doi.org/10.1080/10643380802000974

Kluyver, T., Ragan-Kelley, B., Pérez, F., Granger, B. E., Bussonnier, M., Frederic, J., ... & Willing, C. (2016). Jupyter Notebooks-a publishing format for reproducible computational workflows. *Elpub*, 87-90. 10.3233/978-1-61499-649-1-87

Lehmann, J., Hansel, C.M., Kaiser, C. *et al.* (2020). Persistence of soil organic carbon caused by functional complexity. *Nat. Geosci.* **13**, 529–534 https://doi.org/10.1038/s41561-020-0612-3198718

Robertson, A. D., Paustian, K., Ogle, S., Wallenstein, M. D., Lugato, E., & Cotrufo, M. F. (2019). Unifying soil organic matter formation and persistence frameworks: the MEMS model. *Biogeosciences*, *16*(6), 1225-1248. https://doi.org/10.5194/bg-16-1225-2019

Schmidt, M., Torn, M., Abiven, S. et al .(2011) Persistence of soil organic matter as an ecosystem property. *Nature* 478, 49–56. https://doi.org/10.1038/nature10386

Stegen, J. C., Garayburu-Caruso, V. A., Danczak, R. E., Goldman, A. E., Renteria, L., Torgeson, J. M., and Wells, J. R.: Hyporheic Zone Respiration is Jointly Constrained by Organic Carbon Concentration and Molecular Richness, EGUsphere [preprint], https://doi.org/10.5194/egusphere-2022-613, 2023.

Song, H.S., Stegen, J.C., Graham, E.B., Lee, J.Y., Garayburu-Caruso, V.A., Nelson, W.C., Chen, X., Moulton, J.D. and Scheibe, T.D. (2020). Representing organic matter thermodynamics in biogeochemical reactions via substrate-explicit modeling. *Frontiers in microbiology*, 11, p.531756. https://doi.org/10.3389/fmicb.2020.531756

Stegen, J.C., Johnson, T., Fredrickson, J.K. *et al.* (2018). Influences of organic carbon speciation on hyporheic corridor biogeochemistry and microbial ecology. *Nat Commun* **9**, 585 https://doi.org/10.1038/s41467-018-02922-9

Stephanopoulos, George, Aristos A. Aristidou, and Jens Nielsen. (1998) Metabolic engineering: principles and methodologies.

Tfaily, M. M., Chu, R. K., Toyoda, J., Tolić, N., Robinson, E. W., Paša-Tolić, L., & Hess, N. J. (2017). Sequential extraction protocol for organic matter from soils and sediments using high resolution mass spectrometry. *Analytica chimica acta*, 972, 54-61.

Tolic, N., Liu, Y., Liyu, A., Shen, Y., Tfaily, M. M., Kujawinski, E. B., ... & Hess, N. J. (2017). Formularity: software for automated formula assignment of natural and other organic matter from ultrahigh-resolution mass spectra. *Analytical chemistry*, 89(23), 12659-12665.

Wang, G., W.M. Post and M.A. Mayes (2013). Development of microbial-enzyme-mediated decomposition model parameters through steady-state and dynamics analyses, *Ecological Applications*, 23(1), 255-272, https://doi.org/10.1890/12-0681.1

Ward, N.D., Keil, R.G., Medeiros, P.M., Brito, D.C., Cunha, A.C., Dittmar, T., Yager, P.L., Krusche, A.V., Richey, J.E. (2013) Degradation of terrestrially derived macromolecules in the Amazon River. Nature Geoscience. 6 (7), 530-533. https://doi.org/10.1038/ngeo1817

Ward, N.D, Muller, K.A., Chen, X., Zhao, Q., Chu, R., Cheng, Z., Wietsma, T.W., Kukkadapu, R.K. (2023). Interactive Effects of Salinity, Redox State, Soil type, and Colloidal Size Fractionation on Greenhouse Gas Production in Coastal Wetland Soils. ESS Open Archive. https://doi.org/10.31223/X5FM0N