# 1 NMP-H-CSydro 31.0: a C# Programming Llanguage and

# 2 Windows System based Ecohydrological Model Derived

# 3 from Noah-MP

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- 8 Abstract. The community Noah with multi-parameterization options (Noah-MP) land surface model 9 (LSM) is widely used in studies from uncoupled land surface hydrometeorology and ecohydrology to 10 coupled weather and climate predictions. In this study, we developed NMP-HydroNMH-CS 31.0, a 11 hydrological model written in CSharp(C#). NMP-Hydro NMH-CS 3.0 is a new model was developed by 12 faithfully translating the FORTRAN version Noah-MP from the uncoupled WRF-Hydro 3.0, and wasis 13 coupled with a river routing model. NMP-HydroNMH-CS has the capacity of execution on Windows 14 systems, utilizing the multi-core CPUs commonly available in today's personal computers. The code of 15 NMP-HydroNMH-CS has been tested to ensure that it produces a high-degree of consistency with the 16 output of the original WRF-Hydro. High-resolution (6 km) simulations were conducted and assessed 17 over a grid domain covering the entire Yellow River Basin and the most part of North China. The spatial 18 maps and temporal variations of many state variables simulated by NMP-HydroNMH-CS 13.0 and WRF-19 Hydro/Noah-MP demonstrate highly consistent results, with occasionally minor discrepancies. The river 20 discharge for the Yellow River simulated by the new model withunder various scheme combinations of 21 six Noah MP parameterizations exhibits exhibit general agreement with the natural river discharge at the 22 Lanzhou station. NMP-HydroNMH-CS can be regarded as a reliable replica of Noah-MP in WRF-23 Hydro 3.0, but it can leverage the modern, powerful, and user-friendly features brought by the C-# 24 language to significantly improve the efficiency of the model users and developers.

#### 1.1 Introduction.

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In contemporary hydrological prediction and flood warning applications, the effectiveness of hydrological models hinges on their ability to delineate intricate energy and water processes on the land surface, surpassing the capabilities of traditional rainfall-runoff models. To address this demand, certain land surface models (LSMs) utilized by atmospheric science communities have been bolstered with hydrological simulation features, as observed in WRF-Hydro (Lin et al., 2018) (Lin et al., 2018), or conventional rainfall-runoff models have been enriched with more comprehensive descriptions of land surface processes, exemplified by the VIC model (Liang et al., 1994) (Liang et al., 1994).

The Noah Land Surface Model with multi-parameterizations (Noah-MP) (Niu et al., 2011; Yang et al.,

34 <u>2011)(Niu et al., 2011; Yang et al., 2011)</u>stands out as a robust tool for studying global water issues,

35 serving as the foundation for models like WRF-Hydro, which incorporates Noah-MP (Gochis,

36 2020)(Gochis, 2020). However, the code for Noah-MP and WRF-Hydro is written in FORTARN, a

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37 'legacy' language, posing challenges for code analysis and editing, unlike more modern languages such 38 as CSharp(C#) or Java, because FORTRAN lacks the similar intelligent efficient programming tools that 39 are now common for C#. This limitation makes it arduous for users unfamiliar with FORTRAN to efficiently comprehend and modify the code. Additionally, Noah-MP and WRF-Hydro necessitate a 40 41 UNIX-like operating system, causing inconvenience for the users and developers relying on Windows 42 systems. Therefore, there is a compelling need to code Noah-MP in a contemporary modern 43 programming language, to gain a wider accessibility of the Noah-MP model. 44 We designed NMP-HydroNMH-CS 13.0, a Noah-MP based hHydrological model based on Noah-MP 45 but coded using the CSharp (C#) programming language. This model was crafted by creating a 46 framework and accurately translating the original Noah-MP LSM code from WRF-Hydro 3.0 and 47 coupling with a Muskingum method-based river routing model(Liu et al., 2023)(Liu et al., 2023). C#, Field Code Changed 48 recognized for its modern and object-oriented approach, is widely used for software development across 49 various platforms, particularly on the Windows operating system. <u>According to the TIOBE Programming</u> 50 Community Index (https://www.tiobe.com/) for October 2024, C# ranks fifth among major programming languages with a user base of 5.6%, while Fortran ranks ninth with only 1.8% of users. 51 52 NMP-HydroNMH-CS providesoffers several advantages over the original WRF-Hydro/Noah-MP. 53 Unlike the original version that requires compiling for each computer and primarily relies 54 depends on Unix-like systems, NMP-HydroNMH-CS ean can seamlessly run on Windows systems that 55 supporting the Microsoft Dotnet Framework. The executable files, once compiled, can be easily 56 packaged and distributed to other Windows eomputersmachines, providing offers greater convenience 57 for users who are less familiar with Unix-like operations. \_-The utilization of the C# language facilitates 58 advanced software programstools for code-visualizingation and analyanalyse the model's codesis, 59 enhancing the user convenience for users to eode reading, and modify and debug the code eation. This 60 is appealing to the model developers who are proficient in C# language and the object-oriented 61 programming. The design of modelNMH-CS's design aligns with the input datasets and settings configurations specified in the 'namelist' file, ensuring high compatibility with WRF-Hydro 3.0. Based 62 63 on the support of Leveraging the parallel computing capabilitiesation of C#, both the translated Noah-MP 64 LSM simulation and the river routing simulation with in NMP-HydroNMH-CS support parallel execution 65 on common personal computers.\_ 66 Formatted: Font: (Asian) +Body Asian (等线), (Asian) Chinese (Simplified, Mainland China) 67 2 The Noah-MP LSM 68 Noah-MP is a robust model renowned for its capability to represent diverse physical processes. Since its 69 initial introduction by (Niu et al., 2011) Niu et al. (2011) and (Yang et al., 2011) Yang et al. (2011), Field Code Changed 70 Field Code Changed Noah-MP has been widely used. For example, Noah-MP has been coupled to WRF-Hydro as a major 71 module, and can be has been seamlessly integrated into the Weather Research and Forecasting (WRF) 72 model, giving rise to the WRF-Hydro model, as elucidated by (Gochis, 2020) Gochis (2020). Furthermore, Field Code Changed

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the offline WRF-Hydro model plays a pivotal role in the National Water Model, contributing to the

simulation of floods and river flows across the United States, as highlighted by (Bales, 2019)Bales

(2019), (Francesca et al., 2020) Francesca et al. (2020), and (Karki et al., 2021) Karki et al. (2021). Noah-

MP's versatility extends to applications such as streamflow prediction (Lin et al., 2018)(Lin et al., 2018)

and the estimation of spatial distributions for evapotranspiration, surface temperature, carbon fluxes, heat

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2022; Ma et al., 2017; Yang et al., 2021)(Chang et al., 2020; Gao et al., 2015; Li et al., 2022; Ma et al.,
 2017; Yang et al., 2021). -Noah-MP is supported by several different modelling frameworks to facilitate
 coupling it to various earth system framework models including HRLDAS (Chen et al., 2007),
 LIS(Kumar et al., 2006), and WRF-Hydro(Gochis, 2020). This makes Noah-MP a powerful research and

83 <u>forecasting tool within the hydrology community.</u>

Noah-MP excels in physically representation ofing water and energy dynamics across various—environmental layers, encompassing including a vegetation canopy layer, multiple snow and soil layers, and an optional unconfined aquifer layer for groundwater. Unlike functioning in isolation, Noah MP is typically coupled with host models, such as the community WRF-Hydro modelling framework (Lin et al., 2018) and HRLDAS (Chen et al., 2007), emphasizing its collaborative nature. Additionally, Noah-MP plays a pivotal role in the National Water Model (Francesca et al., 2020; Gochis, 2020), contributing to real-time streamflow forecasts for the entire United States of America. To capture specific physical processes, Noah-MP employs multiple parameterization schemes, offering providing users the flexibility to chooseselect from a total of 12 parametrizations, as detailed outlined in Table 1. This versatility enables tailored representation of diverse environmental conditions and processes, enhancing the model's adaptability and applicability.

Table 1: Parameterization ooptions for Noah-MP (an asterisk (\*) denotes the recommended default option). Certain abbreviations correspond to terms used for Parameterization Schemes (PS) in Noah-MP, and their meanings can be referenced in the Noah-MPWRF-Hydro 3.0 user document (Gochis et. al., 2015). For instance, LAI and FVEG represent the leaf area index and the fraction of vegetation eover, respectively. Note that these options may not be applicable to other versions of Noah-MP, such as that used in HRLDAS. The scheme options here are presented by some abbreviation marks such as 'Noah' or Schaake96 are those used in the 'namelist' file for Noah-MP.

Abbreviation	Physical parameterization	Scheme	Scheme Options
		Code	
DVEG	Vegetation option	1~5	*1 table LAI, read FVEG;
			2 dynamic LAI, FVEG=f(LAI);
			3 table LAI, FVEG=f(LAI);
			4 table LAI, FVEG=maximum;
			5 dynamic LAI, FVEG =maximum
CRS	Stomatal conductance (controls transpiration	1~2	*1 Ball-Berry; 2 Jarvis
	from leaves)		
BTR	$\beta\text{-factor}$ (soil moisture stress factor controlling	1~3	*1 Noah; 2 CLM; 3 SSiB
	transpiration)		
RUN	Runoff (runoff generation at and below the	1~4	1 SIMGM; 2 SIMTOP; *3 Schaake96;
	surface)		4 BATS
SFC	Surface layer drag coefficient	1~2	*1 M-O; 2 Chen97
FRZ	Frozen soil permeability	Fixed to 2	*1 NY06; 2 Koren99
INF	Supercooled liquid water	Fixed to 2	*1 NY06; 2 Koren99
RAD	Radiation transfer option	1~3	1 gap=F(3D,cosz); 2 gap=0; *3 gap=1-
			Fveg

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ALB	Snow surface albedo	Fixed to 2	1 BATS; *2 CLASS
SNF	Precipitation partition option (rainfall or	Fixed to 2	*1 Jordan91; 2 BATS; 3 Noah
	snowfall)		
TBOT	Lower boundary of soil temperature	1~2	*1 zero-flux; 2 Noah
STC	Snow/soil temperature time scheme	Fixed to 1	*1 semi-implicit; 2 fully implicit;
			3 Ts=f(fsno)

## 3 Development of NMP-HydroNMH-CS

#### 3.1 Translation of Noah-MP Code

Our primary focus in developing NMP-HydroNMH-CS involvesd translating the original FORTRAN code of Noah-MP into the C# language. The overarching objective of this translation is to create a hydrological model based on Noah-MP capable of functioning seamlessly on Windows systems. It is essential to note that this translation is based on a relatively older version of Noah-MP utilized in WRF-Hydro 3.0, as the process was started commenced before the release of Noah-MP 5.0 (He et al.,

113 <u>2023</u>)(He et al., 2023).

Converting FORTRAN code into C# is not straightforward due to significant syntax differences insyntax between the two languages. The reconstruction of the model in the C# language follows an straightforward object-oriented design. While FORTRAN is traditionally a function-based language, the core Noah-MP module's functions, subroutines, and state variables are encapsulated as members within a C# class named GridCell (Fig. 1(a)). This class represents all Noah-MP behaviors within a grid box. The variable names, function definitions, data structures, and execution logic have been kept largely consistent with the original FORTRAN code, ensuring user-friendliness for those familiar with Noah-MP. To handle the execution on multiple grid boxes, another C# class named Driver is employed. This class manages tasks such as initializing model variables, creating multiple grid boxes, reading/writing files, and controlling the execution of the model.

Throughout the translation process, a key focus was addressing operations on FORTRAN arrays (**Fig. 1(b)**), crucial for representing the state of soil and snow layers in Noah-MP. Unlike C#, FORTRAN allows arrays to have user-specified index ranges (e.g., index values from -3 to 4). However, in C#, the first index of all arrays invariably starts from 0. To streamline the translation, we **introduced designed** a **newwrapping array** class of C# arrays, named FortArray; \_\_designed to mimic FORTRAN arrays. The wrapped inner array data in FortArray adheres to standard C# conventions, accepting 0 as the inner index of the first element. Yet, externally, the class allow access to the array values through extra indices: by The class providinges methods for index translation from outer indices (FORTRAN style) to inner indices (C# style):

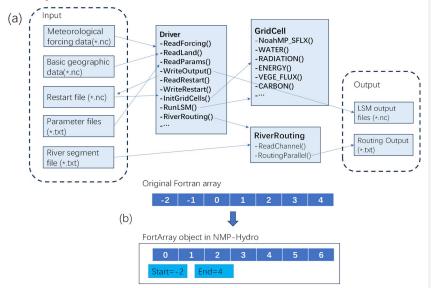
 $135 I_{\rm in} = I_{\rm ex} - I_{\rm start} (1)$ 

Where  $I_{in}$ ,  $I_{ex}$  and  $I_{start}$  represent the inner index, the outer index and the first outer index. The inner index corresponds to the standard C# arrays, while the outer index corresponds to the FORTRAN arrays. For instance, if a FORTRAN array of 8 elements has an index range from -3 to 4, this arrayit will be is translated into a FortArray that has a standard inner array of 8 elements, and accompanied by two

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arguments representing the starting FORTRAN index (-3) and the ending FORTRAN index (4), but then the range of its inner indices are remain 0~7. This array translation technique ensures that all the original execution logic in Noah-MP is seamlessly preserved in NMP-HydroNMH-CS.

The model also supports parallel execution, implemented through the native parallel functionality of the C# language. The technique canse functions efficiently allocate computational tasks for distinctover the grid boxes to different CPU threads which can be executed by separate CPU cores. For instance, if a specific grid domain requires the execution overof 2400 grid boxes, and the tasks are assigned to 8 threads, each thread is responsible for completing the tasks calculations one approximately 300 grid boxes. It's crucial to note that if the number of specified threads exceeds the actual number of CPU cores, severalmultiple threads should be may end up executed bying on the same single CPU core. Therefore, specifying more threads than the available CPU cores does not contribute to an overall improvement in execution speed.



**Figure 1.** The architectural diagram of NMP-Hydro NMH-CS (a) and the conversion of FORTRAN arrays to C# arrays (b). NMP-Hydro is a reconstructed replica of the version of Noah-MP that is coupled in WRF-Hydro 3.0.

## 3.2 Coupling with a parallel river routing module

The Noah-MP land surface model can produce column outputs of runoff, but cannot simulate the horizontal movement of water. In order to simulating the surface movement of runoff in the river channels, we In the development of NMP-Hydro, we integrated a parallel river routing module, which is —based on the Muskingum method for vectorized channel networks. methodThe module is described in a previous study (Liu et al., 2023)<sub>72</sub>. This module is deviatingnot —from—the previous utilization of the eoupled-RAPID model that was coupled in WRF-Hydro (Lin et al., 2018). This parallel river routing module, implemented using C#, incorporates our unique techniques.:—

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The first technique is aAn array-based sequential processing method for Muskingum routing.  Muskingum Curves agustion (Curves 1060) with letteral inflavor appridered in	Formatted:	No widow/orphan control	
Muskingum-Cunge equation (Cunge, 1969) with lateral inflow considered is	E		
$Q_{e,t+1} = C_0 * Q_{s,t} + C_1 * Q_{s,t+1} + C_2 * Q_{e,t} + C_3 * Q_{lat,t+1} $ (2) Where	Formatted:	-ont: 10 pt	
$C_0 = \frac{kx + .5\Delta t}{k(1-x) + 0.5\Delta t},  C_1 = \frac{-kx + 0.5\Delta t}{k(1-x) + 0.5\Delta t},  C_2 = 1 - C_0 - C_1,  C_3 = \frac{\Delta t}{k(1-x) + 0.5\Delta t}.$	Formatted		
Here, O represents the channel streamflow (m <sup>3</sup> /s), which can be considered as a function of time and	Formatted Formatted		
position. s is the start point of a channel segment/node, e denotes the end point of the channel segment,	Formatted		
both of them are used as subscriptions for different spatial positions. <i>n</i> denotes the start of the period or	Tormatted		(
inflow, $n+1$ denotes the end of the period or outflow. Both $n$ and $n+1$ are used as subscriptions for position.			
The subscription 'lat' represents the lateral streamflow(runoff) from current river catchment.			
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Two parameters of the Muskingum-Cunge method are $k$ and $x$ . $k$ is the travel time of a flow wave with		ndent: First line: 0 cm	
celerity $c_k$ through a channel segment of length $L$ , thus $k=L/c_k$ . Parameter $x$ can be estimated by			
$x = \frac{1}{2} \left( 1 - \frac{q}{S_0 c_b L} \right) \tag{3}$			
Where $q$ represents unit width streamflow, $S_o$ is the channel bed slope.	Formatted:	Font: 10 pt	
The lateral inflow of each river segment is the runoff simulated by column Noah-MP LSM in NMH-CS,	Formatted:	ndent: First line: 0 cm	
which is expressed as $Q_{\text{lat},t} = R_{c_{s,t}} A_{s}$ (4)			
Where $Q_{\text{lat},t}$ is the lateral inflow at time $f_{\text{e}}$ $R_c$ is the runoff value at the given grid box, $A_s$ is the local	Formatted		
catchment area of current channel segment.	ronnatteu		<u> </u>
catelinien area of current channel segment.			
The form of Eq.(2) implies that the river routing calculation can be completed from any upstream river	Formatted		
segment to its downstream segment. At the time $L^{+1}$ , all the outflows $(Q_{e, t+1})$ of multiple upstream	Formatteu		(
segments will be summed up as the inflow $(Q_{2,4+1})$ of current segment. Although a common river network	'		
is a tree-like structure, it can be represented as a sequential array, where any upstream river segment is			
stored near the array head (with zero index), while its downstream river segment is stored near the array's			
tail (with large array index). Therefore, the Muskingum routing can be calculated over the river segment			
array in a unidirectional processing way (please see Fig.1 in Liu et al. (2023)).			
array in a uniqueculonar processing way (please see Fig. 1 in pla et al. (2023)).	Formatted:	Font: 10 nt	
(1)—The second technique is the		Normal, No bullets or numbering	
A straightforward equal-sized domain decomposition method, to conduct parallel calculation: just	Formatted	vormal, two bullets of numbering	
allocating the river segments into equal-sized blocks. Within each block, the Muskingum routing can be	Tormattea		(
executed separately by a single CPU core. This treatment is based on that in any block, most segments			
have upstream segments within the same block, and only a small fraction of the segments have upstream			
segments in other blocks. Therefore, all river segments that receive inflows from other blocks (referred			
to as 'cross-block segments') need to be identified. These cross-block segments should be executed by			
the primary core, after the multi-domain parallel execution is completed.:			
(2)	Formatted:	Normal, No bullets or numbering	
(3) The third technique is a Three distinct parallelization schemes for river routing.		ndent: Left: 0 cm, First line: 0 cm	
4) A specific sorting approach for river segments used in domain decomposition. It has been proven that	Formatted	nache Leit. O Gill, Filst IIIIe. O GIII	
a depth-first traverse of the river segments is more suitable for the parallel execution of the Muskingum	Tomatted		
a department develop of the river segments is more suitable for the parameter execution of the Muskinguin //			

This module requires two additional inputs files, a river segment list file named 'ChannelOrder.txt' and a 'namelist.txt' file. The latter file is used to set parameters and the length of time step. Each river segment in the list file presents following information: its own index, the index of its next downstream river segment, the row number and the column number of the grid box (in Noah-MP's running domain) providing runoff input to the current segment, the length (m) of the current river segment, the two parameter values (K and X) of the Muskingum method, the area of the catchment of the current segment.

The river segment list can be derived from both gridded river network or vectorized river network. The resolution of the river routing is determined by the original river network from which river segments is derived. Therefore, the choice of using vector river network or gridded river network and the selection of spatial resolution are completely determined by the users. The length of the temporal step of the river routing is required to be multiple times shorter than the time step for running the Noah-MP, and can also be designated by the users. For example, the time step of routing is set to 600s, while the time step for Noah-MP LSM is usually set to 3 hours.

This approach's primary advantage lies in its ability to straightforwardly simply decompose any river network into multiple domains with an equal number of river segments. Achieved by evenly dividing the river segment list into any number of blocks, this innovation capitalizes on the inherent tree-like structure present in most river networks. Importantly, it does not necessitate consideration of the topological conditions specific to a given network, as required in studies such as Mizukami et al. (2021) or David et al. (2015). This design allows parallel execution of river routing on modern personal computers equipped with multi-CPU cores.

The integration of the river routing module with the Noah-MP LSM involves assigning lateral inflows from the LSM-simulated total runoff to the river routing model. In the present NMP-HydroNMH-CS configuration, we utilize a straightforward-catchment centroid-based coupling interface (David et al., 2015). This method designates the LSM grid cell containing the catchment centroid (referred to as the "centroid cell") as the location for a river reach to receive lateral inflows. At a specific temporal step, the computed contributing runoff discharge  $Q_{lat}$  (unit: m³/s) is determined by the following expression:

 $Q_{lat} = R(nx, ny) \times F \times 1000$  (5)

where R(nx, ny) is the runoff (mm, surface + subsurface) simulated by the LSM during the time step, F is the catchment area (km<sup>2</sup>) contributing water to the current river segment.

Alternatively, employing weighted assignments from different grid boxes, akin to the method utilized in (Lin et al., 2018), is also a valid approach. However, this method requires the generation of weights from multiple grid boxes. Given the <u>rough resolution of the meteorological datasetssize of each grid box</u> (equivalent to the resolution of meteorological data, typically ranging from 25 km to 100 km), and <u>considering that</u> each grid box can encompass the catchment areas of multiple river segments, the coupling approach using area weighting is unlikely to yield substantial improvements for most river segments.

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#### 3.3 Code Debugging Process

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281 282 To eliminate any potential code errors resulting from incorrect translation, we conducted a thorough checking review of the code by performing model execution benchmark tests on single-column running on specific grid boxesa grid domain. Here, In the large domain (the same domain described in section 4.1), each time the grid box for single-column execution is arbitrarily selected. Such debugging These tests were carried out in twothree approaches stages. The first approach was Initially, the code underwent carrying out a meticulous step-by-step eheckdebugging by examining the printed values of manyeach variables (including many local variables in the code) in WRF-Hydro 3.0-running for specific single grid boxes. \_-This debugging-process was also conducted repeated by switchingon each option of multiple physical parameterization schemes. The grid box for the single-column debugging was also switched several times. Such debuggingThis process has been conducted numerous times and has effectively eliminated any code errors arising from inaccurate translation. Although meteorological driving data for the debugging simulation is prepared for the period between 2000 and 2016, However, it is important to note that such debugging these tests are only feasible for a limited number of temporal steps-andfor a -grid-box execution. It is also impossible to conduct debugging on the entire domaines. The second approach is an artificial code checking process. Considering that the stepwise <u>Ddebugging</u> the code through years-long simulations across the entire domain is impractical, we checked the NMH-CS's code by comparing with the original FORTRAN code for many times. Through this checking, many code inconsistences were identified and corrected. -In the second phase, we compared the spatial distribution of multiple state variables and the long-term

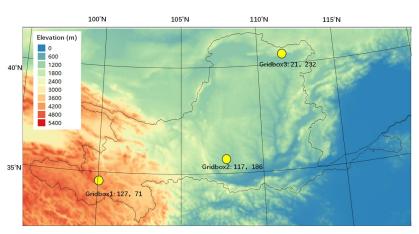
In the second phase, we compared the spatial distribution of multiple state variables and the long term discharge hydrograph produced at the main river sections by NMP-Hydro with that produced by WRF-Hydro 3.0. After identifying and correcting any erroneous code, the final results indicated that the two models are capable of generating broadly identical outcomes.—

# 274 4. Benchmark Testing of NMP-HydroNMH-CS

#### 4.1 Application area and data

#### 4.1.1 Application area

The Yellow River Basin in Northern China is used as a test area of NMP-HydroNMH-CS. The gridded domain, as illustrated in Fig. 22 and Fig. 33, encompasses the entirety of the Yellow River Basin (YRB) and the most part of North China, comprisingnsisting of 350 columns and 170 rows, featuring with a resolution of 6 km in Lambert conformal conic projection coordinates. Geophysical information data essential for the domain, including digital elevation, land use and land cover, and green vegetation fraction, wereas extracted from the WRF/WPS 3.5 input database.



**Figure 2.** The terrain map of the simulation domain and the three grid boxes for comparison of state variable time series. The three grid boxes <u>for extracting state variables to compare</u> are represented by yellow dots, marked with grid box code, the row number and the column rows.

For the river routing simulation, the digital river network of the Yellow River was obtained from the HydroSHEDS dataset (version 1) (http://hydrosheds.cr.usgs.gov/). HydroSHEDS wasis derived from gridded digital elevation data with a resolution of 15 arc-seconds. Given the substantial human intervention and the intricate nature of reproducing observed daily or hourly water discharge, uniform values were assigned to all river segments for the river routing parameters (specifically, the wave celerity ( $c_k$ ) and another parameter (x) describing the river channel condition, as detailed in (David et al. ( $c_1$ 2013)). No precise calibration is required here, as monthly or annual river discharge remains unaffected by changes in routing parameters.

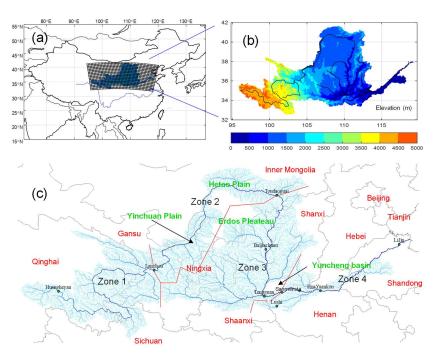
The Yellow River basin experiences significant human impacts, including irrigation, industrial water usage, and groundwater extraction. Major artificial reservoirs and numerous smaller reservoirs regulate the river's discharge, serving as the primary water resource during the dry season. However, such extensive human interference presents substantial challenges in accurately modeling river discharge. Comparatively, the river discharge upstream of the Lanzhou (In Zone 1 as shown Fig.23) hydrological station contributes over half of the entire YRB's total discharge, and is relatively less impacted by dams, enabling us to objectively test the model's performance.

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**Figure 3.** The grid domain covering the Yellow River Basin and the North China area: (a) Geographical location within China; (b) Elevations; (c) Vector River networks utilized for river routing modeling (extracted from the HydroSHEDS dataset). The delineation of the boundaries between distinct zones controlled by four gauging stations (Lanzhou, Toudaoguai, Sanmenxia, Lijin) is represented by red lines.

#### 4.1.2 Meteorological dataset and river discharge data

To drive the two models (NMH-CS and WRF-Hydro), the samea 3-hourly and 6-km grided meteorological forcing dataset compriseding of shortwave and longwave downward radiation, wind velocity, air temperature, relative humidity, air pressure at the surface, and precipitation rate was acquired. The benchmark dataset was extracted clipped and regrided (bilinear interpolation)—from the 1.0°×1.0° GLDAS-1 land surface product (Gan et al., 2019; Rodell et al., 2004)(Gan et al., 2019; Rodell et al., 2004), for the period 2000-2016. Given the limited availability of observational/reconstructed\_river discharge data between 2001 and 2016, the extracted data pertains to this period, and additional data between 1996 and 2000 was also extracted for the model's spinning-up.

In previous research, the spinning-up of Noah-MP requires 50 years (Wu et al., 2021)(Wu et al., 2021) or more than one hundred years (Zheng et al., 2019) (Zheng et al., 2019) to achieve an equilibrium state. However, in this study, the spin-up process was conducted in two steps. In the first step, the period from 1996 to 2016 was run three times to generate a 'restart file' for a 63-year spinning-up, utilizing the initial PS combination (11131-2222-121). In the second step, starting from this initial combination, new schemes were introducedadopted, and the 'restart file' obtained from the initial scheme combination was used to initiate the formal experiments covering the period from 1996 to 2016.

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The dataset of Natural River Discharge (RND) reconstructed by the Yellow River Conservancy Commission of the Ministry of Water Resources was gathered to assess the model output. Annual natural discharges from the monitoring station of Lanzhou were collected for the period from 2001 to 2016.

#### **4.2 Running speed**

Compared to other differences between the two models, running speed is the least important factor to consider. Considering that NMH-CS and WRF-Hydro run on different platforms (Windows or Linux) and machines, it is also difficult to achieve a comparative evaluation of running speed. Actually, the comparison of running speed depends on the programming language used. In theory, FORTRAN and C programs can run faster than C# programs because FORTRAN and C are relatively low-level languages compared to any modern object-oriented language. However, as a language that can run in native machine code, C# is not slow. This means that there won't be a large difference in running speed between C# and FORTRAN. There are few authoritative publications on the running speed of these two languages, but there have been many documents on benchmark testing on the internet. When considering parallel execution, comparing the running speed of the two models also become more unnecessary. NMH-CS can run in parallel mode on personal computers, while WRF-Hydro does not have this functionality. On the other hand, WRF-Hydro can run in parallel mode in the Message Passing Interface (MPI) environment of high-performance computers, while NMH-CS does not support MPI.

We tested the execution time of NMH-CS by setting different numbers of C# parallel threads. The computer used for the testing is a common laptop with 6 CPU cores. The results indicate that for the execution of the entire domain, as the number of threads increases from 1 to 6, the average time consumed per time step is 1576ms, 977ms, 801ms, 711ms, 679ms, and 672ms, respectively. When the number of threads is set to 1, the time spent is slightly greater than the execution time in the non-parallel mode (1461ms). It is worth noting that the time spent is not linearly related to the number of parallel threads, which can be explained by various reasons. One is that some tasks are not actually executed in parallel mode, such as reading meteorological input files. Another reason is that not all threads in NMH-CS are fully processed by the CPU cores, as there are many other tasks in the entire Windows environment that have to be processed simultaneously by the CPU cores.

#### 4.23 Comparing the outputs of NMP-HydroNMH-CS and WRF-Hydro3.0

It is noteworthy that there are numerous parameterization scheme combinations for Noah-MP, which makes it unfeasible to compare the results generated under all scheme combinations. Therefore, the output of NMP-HydroNMH-CS and WRF-Hydro was compared only with under—the default parameterization scheme combination was compared, based on the exact same meteorological dataset. The comparison was conducted in two ways. The first comparison is that of the spatial maps of multiple variables (Table 2) for a specific year or day. For each state variable, such as SFCRNOFF, the maps of state variables simulated by NMP-HydroNMH-CS and WRF-Hydro are presented. The difference ( $\triangle$ ) between the values of the same variables simulated by the two 'Noah-MP models' is calculated as  $\triangle = V_{NMP-HydroNMH-CS} - V_{WRF-Hydro}$  (6)

Where  $V_{\text{NMP-Hydro}\text{NMH-CS}}$  and  $V_{\text{WRF-Hydro}}$  are the state variable simulated by the two Noah-MP models, respectively.

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**Formatted:** Font: (Asian) +Body Asian (等线), (Asian) Chinese (Simplified, Mainland China) 366 For certain variables, such as SFCRNOFF, UGDRNOFF, ECAN and ETRAN, the percentive relative 367 differences were also calculated as follows:

368  $\delta = 100 \cdot \Delta / V_{\text{WRF-Hydro}}$ 

The second method is to compare the <u>temporal</u> variations of state variables at specific grid boxes. In this case, only three grid boxes (Fig. 2) were selected to extract the state variable time series. The selection of these grid points is an arbitrary decision made by roughly considering different climate zones, without other strict consideration. Gridbox1 is selected from the Qinghai Plateau region, corresponding to the source region of the Yellow River. Gridbox2 corresponds to a location of Inner Mongolia and the north of Loess Plateau. Gridbox3 is in a hilly area of the Wei River Basin (a major part of the Yellow River Basin). In fact, during this study, other grid points also have been casually tested, but the results are mostly similar to the above mentioned 3 grid boxes, and will not be presented in the paper.

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Table 2 The state variables simulated by NMP-HydroNMH-CS and WRF-Hydro, which is verified by generating maps

by generating maps		
Variable name	description	unit
SFCRNOFF	Accumulated Surface runoff	mm
UGDRNOFF	Accumulated ground runoff	mm
ECAN	Evaporation from canopy	mm
ETRAN	Vegetation transpiration	mm
TV	Vegetation temperature	K
TG	Ground temperature	K
SOILT	The temperature for soil layers	<u>K</u>
SOILW (or SH2O)	The volumetric content of moisture in soil layers	<u>m³⋅m⁻³</u>
<u>SNOWH</u>	The total depth of snow layer	<u>m</u>
SNEQV	Snow water equivalent	kg·m <sup>-2</sup>
EAH	Canopy air vapor pressure	Pa
EVG	Ground evaporation heat	W <u>-</u> /m <u>-</u> 2
CHV	Exchange coefficient vegetated	m <u>.</u> /s <u>-1</u>
CHLEAF	leaf exchange coefficient	
TR	Transpiration heat	$\underline{\text{W}\cdot\text{m}^{-2}}\underline{\text{W}/\text{m}^{2}}$
EVB	Evaporation heat to atmosphere bare	$\underline{W \cdot m^{-2}} \underline{W/m^2}$
FIRA	Total net long-wave radiation to atmosphere	$\underline{W \cdot m^{-2}} \underline{W/m^2}$
TRAD	Surface radiative temperature	K

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### 4.32.1 Maps of state variables simulated

To test whether NMP-HydroNMH-CS can produce the corresponding outputs of the original WRF-Hydro (Fortran-version Noah-MP), many state variables (Table 23) from multiple time slices (10 June 2000, 10 June 2001, 10 June 2004, and 10 June 2008) have been were checked by drawing maps. - Only four slices (10 June 2000, 10 June 2001, 10 June 2004, and 10 June 2008) were arbitrarily selected here without special consideration. Only some maps of these state variables at certain time slices are presented in both the paper and the supplementary information. The maps for all the sate variables in Table 2 reflect

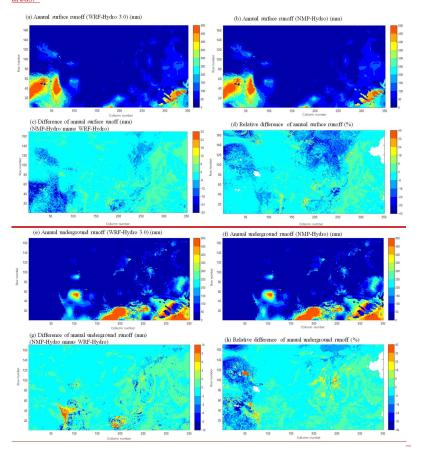
Surface albedo

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Formatted: Superscript Formatted: Superscript high consistence between NMH-CS and WRF-Hydro, with only Tthe maps for two representative variables (SFCRNOFF and TV) are shown in Fig.4 and Fig.65. As can be seen, there is visually little difference in the spatial patterns of the results. Similarly, no discernable visual difference is also apparent for the maps of other variables. However, the relative difference of annual surface runoff and annual underground runoff is significantly large at some areas (generally in high-elevation regions), where NMP-HydroNMH-CS underestimated those values by above 10%-30% (Fig.4). For most of the domain Generally, the difference in TV is smaller than 0.2 °C, but in some sporadically districted locations, the TV's difference can be larger than 2 °C (Fig.56). The comparison of TG has the similar effects, but the difference is more significant than that of TV. The similar high consistency effects are also reflected by other state variables, including soil temperature, soil water content, snow water equivalent (Fig.S2-S4 in supplementary information). The differences in these state variables between the two models are generally small, except some large ones sporadically distributed in the high-latitude



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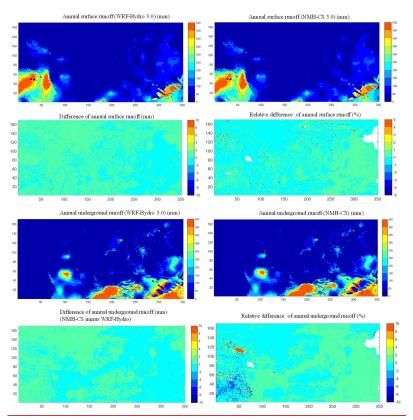
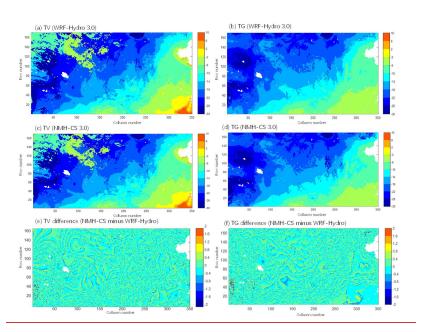
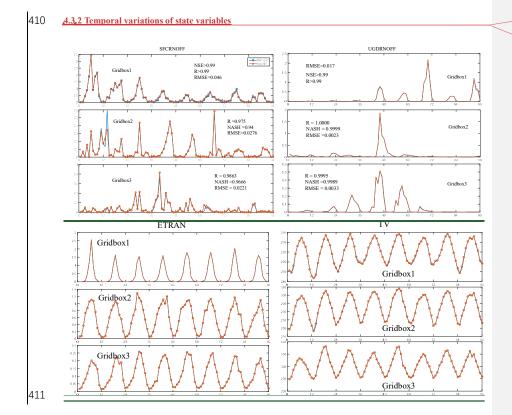


Figure 4. Maps of annual total values, differences, and relative differences of SFCRNOFF (surface runoff, mm) and UGDRNOFF (underground runoff, mm) simulated by WRF-Hydro3.0 and NMP-HydroNMH-CS +3.0, for the year 2005.



**Figure 5:** Maps of TV (vegetation temperature, °C) and TG (ground temperature, °C) simulated by WRF-Hydro3.0 and NMH-CS 3.0, for the day Jan. 1<sup>st</sup>, 2008.



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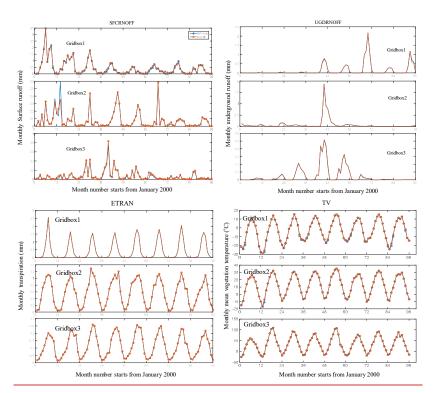


Figure 68: Monthly surface runoff (SFCRNOFF in mm), underground runoff (UGDRNOFF in mm), transpiration (ETRAN in mm) and vegetation temperature (TV in °CK) simulated by WRF-Hydro3.0(Noah-MP, blue) and NMP-HydroNMH-CS (red) at the three grid boxes, 2000-2007.

The monthly temporal variations outputs at the of three representative grid boxes (as shown in Fig.2) indicate that the two models produced consistent temporal changes (Fig.86, Fig.85 and Fig.86). For certain variables, for example, SFCRNOFF—, TV and TG (other variables as well), FIRA, CHV, occasionally, some significant differences wereas found in certain months for Gridbox2. It must be such occasionally happened differences that caused the spatial disparity as shown in Fig.4-5. We checked the disparity at certain grid boxes on the 3-hourly daily sealevalues and found that the differences also happen sporadically (Fig.57). Most of Almost all the disparities fferences occur during the cold months (November, December, January, and February). However, it is worth noting that mostly—of the simulated state variables SFCRNOFF in these months show no difference—and this—difference is also independent of whether precipitation occurred during these days. Each difference is caused by a mismatch in the simulated peak time (Fig.5), manifested as a one day lag effect of NMP—Hydro relative to WRF—Hydro. Meanwhile, this lag effect of NMP—Hydro also causes underestimated or overestimated total surface runoff. Considering such mismatch usually happens in cold months and high-elevation regions, it may be caused by the different calculation accuracy for the processes of ice, snow orand frozen soil. For the three representative grid boxes, no significant differences were

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identified for certain variables, including many variables like TR, EAH, TV, ETRAN, UGDRNOFF (no plots for these variables will be presented in the paper).

By comparing and analyzing the printed state variables (in 3-hourly timesteps) in WRF-Hydro and the NMH-CS, we found the major inconsistencies occur in the module of snow water (named 'SNOWWATER' in the code). —Although the code for these processes in Noah-MP is complex enough, we have checked the code multiple times and have not found any coding difference, so these differences are likely the result of floating-point errors.

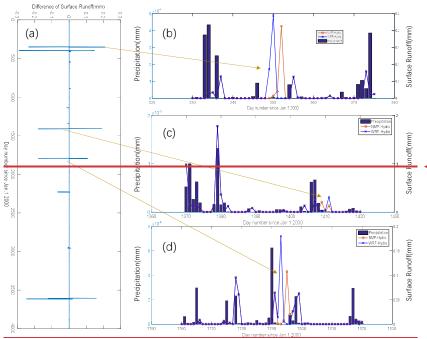


Figure 5. Simulated difference in daily surface runoff (SFCRNOFF) series simulated by NMP Hydro and WRF-Hydro for Gridbox3. (a) the SFCRNOFF difference of NMP Hydro to WRF-Hydro; (b-c) the disparities of SFCRNOFF near the day number 350, 1410 and 1790, respectively. From Fig. 7, three major inconsistencies simulated between the two models occur usually simultaneously in the multiple state variables. These three cases demonstrates that almost all the major inconsistencies in multiple variables are caused by the minor inconsistencies in SNOWH (the state variable to indicate the depth of snow). The logic of snow process in Noah-MP is coded as when SNOWH is below 0.025m, the ISNOW (a state variable to indicate whether snow a layer exists) is set to zero (no snow layer), otherwise, is set to 1 (having a snow layer). Therefore, if SNOWH simulated by NMH-CS is close to 0.025, a small floating-point error may trigger a division between having a snow layer and no snow. Due to the different physical effects of radiation balance between snow layers and the ground, the distinction between having a snow layer and no snow layer will further lead to significant inconsistencies in snow depth (SNOWH), snow water equivalent (SNEQV), soil water (SOILW), vegetation temperature (TV), and ground temperature (TG). Once an inconsistency occurs, it will persist for a period of time. It is highly probable that the minor differences in SNOWH is caused by

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accumulation of floating-point error, because for most of the times the differences are very small except those during the inconsistent periods.

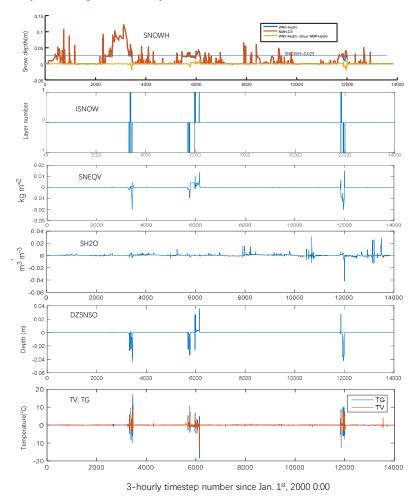


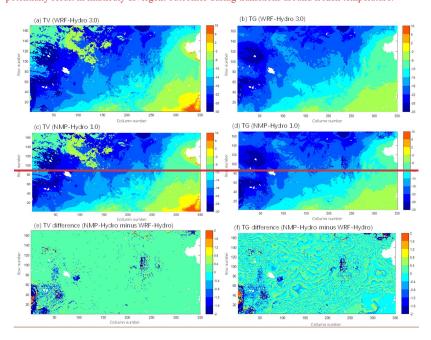
Figure 7: The differences (NMH-CS minus WRF-Hydro) between the three-hourly variables simulated by NMH-CS and those simulated by WRF-Hydro. SNOWH: snow depth (m); ISNOW: Number of snow layers, count; SNEQV: Snow water equivalent (kg·m²); SH2O: soil liquid water content (m³·m³), equivalent to SOILW; DZSNSO: snow/soil layer depth (m). These variable names are those used in the programming code. The occurring of the three inconsistencies correspond to the short periods: March 4, 2001, January 15, 2002 to February 8, 2002, and January 22, 2004 to February 9, 2004.

Given these discrepancies on a monthly scale, no NSE smaller than 0.9 and correlations smaller than 0.98 were identified. For the three representative grid boxes, no significant differences were identified for certain variables, including TR, EAH, TV, ACCETRAN, UGDRNOFF, and others (Fig.8). The daily

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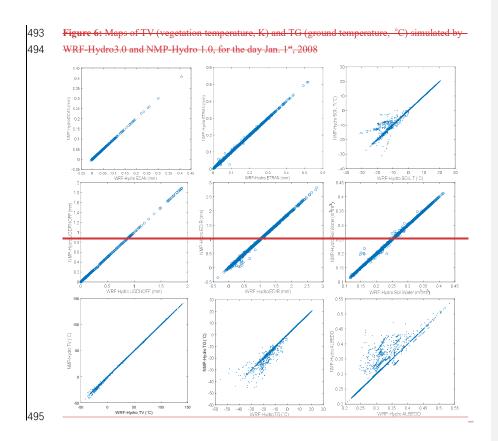
time series from multiple Gridbox3-grid boxes (including the three in Fig.2) were are extracted, then and are compared between the NMP-HydroNMH-CS and the WRF-Hydro. Similar effects were obtained for the grid boxes, but only the results for Gridbox3 are presented as representative in Fig.8.— It is evident that EDIR, SFCRNOFF, soil water content (SOILW) and TV exhibit smaller discrepancies, whereas TG and ALBEDO-demonstrate larger disparities (Fig.7). The discrepancies are pronounced when TG and TV below zero. The daily samples for soil temperature, soil water content, snow depth and snow water equivalent are presented in Fig.S1, Fig.S5 and Fig.S6 (supplementary material). The comparisons about soil layers reflect that the soil temperature has relatively large inconsistencies, which should also be explained by the different division of snow layers that is caused by error when SNOWH approaches 0.025m.

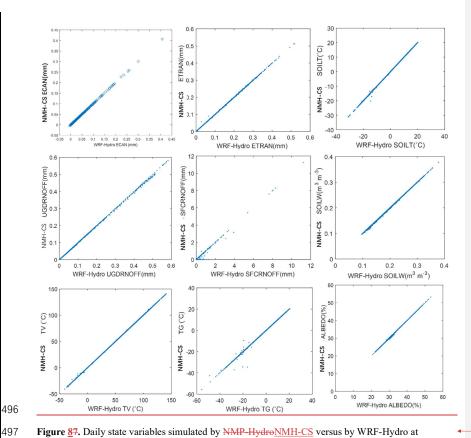
The reason for occasional discrepancies between the outputs of two models remains elusive, as it is challenging to ascertain through the process of code checking and debugging. Such discrepancies may be attributed to a number of factors, including floating point calculation errors, disparate parameter configurations, or encoding inconsistencies. The former two are reasonable and acceptable, whereas a coding mismatch is typically unacceptable. Nevertheless, identifying discrepancies is only feasible during the initial stages of debugging, but not for tens to hundreds of subsequent iterations. It is not uncommon for errors to remain undetected even after the execution of numerous time steps. In this study, in light of the fact that no inconsistencies remained after checking the code many times, it is plausible that floating point errors play a significant role in explaining the discrepancies. During the debugging process, we found inconsistency always arise from the recursive calculation of energy transferring for vegetated and bare land, such as variables TV and TG. The floating point error of temperature could potentially result in markedly divergent outcomes during transitions around frozen temperature.



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**Figure <u>87.</u>** Daily state variables simulated by <u>NMP-Hydro NMH-CS</u> versus by WRF-Hydro at Gridbox3. Due to the high consistency for most of the values, statistical evaluation metrics such as correlation coefficients or relative biases will not be presented in the paper.

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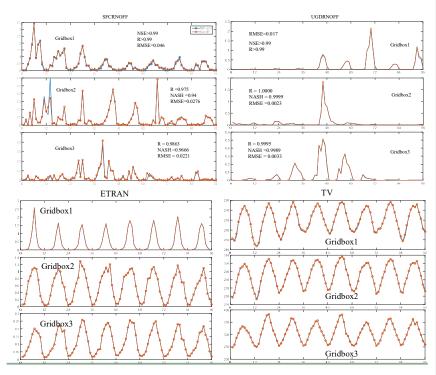


Figure 8: Monthly surface runoff (mm), underground runoff(mm), transpiration (mm) and vegetation-temperature (K) simulated by WRF-Hydro3.0(Noah-MP, blue) and NMP-Hydro (red) at the three grid-boxes, 2000-2007.

# $\textbf{4.3\underline{4} Streamflow discharges } \underline{\textbf{simulated by multiple experiments}} \underline{\textbf{for the Yellow River by NMH-CS}}$

# 4.34.1 Experimental design of Noah-MP simulation

Here, we present the numerical outputs of NMP-HydroNMH-CS and the comparison with WRF-Hydro/Noah-MP on the streamflow discharges over the Yellow River, with various parameterization schemes used. The parallel speedup of NMP-Hydro will not be evaluated here, as it is a straightforward implementation in C# for executing multiple tasks. The description of parallelization of the coupled river routing models has been clearly described in our previous publication(Liu et al., 2023).

To verify whether the various parameterization schemes (PSs) of NMP-HydroNMH-CS can produce reasonable discharge for the Yellow River catchment area, this study conducted 17 Noah-MP simulations using different PS combinations. Given the challenge of determining the relative importance of each parameterization and the impracticality of including all possible combinations, we adopted a strategic approach. A fixed PS combination was established as a foundation, and alterations were made to one parameterization's scheme at a time (refer to Table 3).

In addition to our selected parameterizations, we considered commonly used PS combinations,

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including the 'default' combination proposed by Noah-MP developers. Sensitivity analysis was conducted by analyzing the differences or variations among these incomplete PS combinations. It is important to note that the chosen PS combinations represent only a subset of all possible combinations, and the assumed sensitivities based on this subset are considered indicative of overall sensitivities based on the complete set of combinations.

The PS combinations are represented by codes consisting of sequential digital numbers. For instance, the default combination is denoted as '11131-1132-111', where each number signifies a scheme option. The initial experiment, arbitrarily set as the PS combination of '11131-2222-121', served as the foundation for subsequent experiments. Fifteen experiments (refer to Table 2) were then conducted by modifying one option at a time from the initial experiment.

These experiments are categorized into multiple groups, with the initial experiment '11131-2222-121' being employed in multiple groups:

Runoff scheme group (four experiments, switching between: 1. SIMGM, 2. SIMTOP, 3. Schaake96, 4. BATS):

Vegetation scheme group (five experiments, switching between the first option and the fifth option, see Table 1);

β-factor option group (three experiments, switching between Noah, CLM, and SSiB);

Radiation transfer option group (three experiments, switching between three options);

Group for the scheme of the lower boundary of soil temperature (six experiments);

Group for stomatal conductance scheme (two experiments, switching between two options).

Table 3. Experiments conducted in this study

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Number	PS	combination	Abbreviated	Description
	code		code	
1	11131	-2222-121	11131	The control first experiment
			or 11131-222	
			or 11131*121	
2	11111	-2222-121	11111	Experiments with RUN
3	11121	-2222-121	11121	
4	11141	-2222-121	11141	
5	21131	-2222-121	21131	Experiments with DVEG
			or 21131*121	
6	31131	-2222-121	31131	
7	41131	-2222-121	41131	
8	51131	-2222-121	51131	
			or 51131*121	
9	11231	-2222-121	11231	Experiments with BTR
10	11331	-2222-121	11331	
11	11131	-2212-121	11131-221	Experiments with RAD
12	11131	-2232-121	11131-223	
13	12131	-2222-121	12131	Experiments with CRS
14	11131	-1132-111	'default'	The default PS combination proposed by Noah-MF
				developers
15	11131	-2222-111	11131*111	Experiments with TBOT

16	21131-2222-111	21131*111
17	51131-2222-111	51131*111

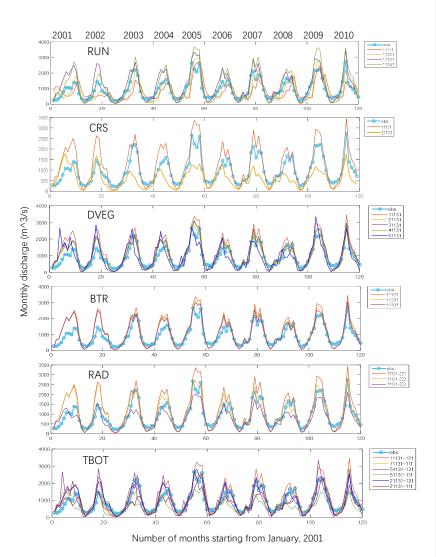
#### 4.34.2 Simulated streamflow under various parameterization schemes

We simply use <u>tThe Taylor diagrams (Taylor, 2001)(Taylor, 2001)</u> are used to evaluate the different PS on the river discharge at the Lanzhou station. Taylor diagram provides a graphical representation of a model's simulation performance, encompassing three key indices: correlation coefficient (*R*), root-mean-square error (RMSE), and standard deviation (SD).

The streamflow discharges were produced by coupling the NMP-HydroNMH-CS with the parallel river routing model. The results showed that, in general, the various scheme combinations of parameterizations can produce monthly river discharge close to the result of the original Noah-MP (WRF-Hydro) for the Lanzhou station. A preliminary comparison on the various scheme combinations is presented in Table 4. The monthly performance is summarized in Table 4 based on the comparison of the different PSs as shown in Fig. 9. It can be observed that for the majority of parameterizations, the discharges in winter are not sensitive to the schemes, this is to be expected, given the minimal runoff during this season. The simulated summer discharges exhibit notable degree of sensitivity with regard to the various parametrization schemes. In relation to the runoff parametrization, the results obtained through the utilization of the SIGGM scheme led to overestimation during the winter season and underestimation during summer, signifying that considering groundwater could enhance the simulation accuracy of the catchment modulation as opposed to other schemes.

For the Lanzhou station, over 50% of experiments produced discharges with correlations larger than 0.9 (Fig.10). The PS combination '11141-2222-121' yielded the highest correlation, and '11131-2222-111' showed the highest performance according to Taylor's score.

Field Code Changed



**Figure 9:** Simulated Mmonthly relief iver discharge (m³/s) for Lanzhou by NMH-CS. The first subplot displays the results simulated with varying RUN schemes while the other subplots follow a similar pattern. Reconstructed natural discharge is denoted as 'obs'.

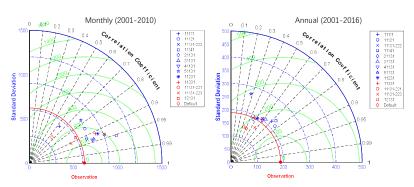


Figure 10: Taylor diagram for monthly and annual mean river discharge  $(m^3/s)$  at the Lanzhou monitoring station

Table 4: Performance of various parameterization schemes on monthly discharge for Lanzhou station

	scheme	Winter	summer
RUN	1.SIGGM	Overestimation	underestimation
	2.SIMTOP		Small overestimation
	3.Schaake96	underestimation	Large overestimation
	4.BATS		The Largest overestimation
CRS	1. Ball-Berry	Small difference	Overestimation
	2. Jarvis	Small difference	Underestimation
DVEG	1.Table LAI, read FVEG		The largest overestimation
	2.dynamical LAI and FVEG=f(LAI)	No significant	Mostly small
	3. table LAI, FVEG=f(LAI)	difference	overestimation
	4 table LAI, FVEG=maximum		
	5.Dynamical LAI, maximum FVEG	Unstable overestimation	on and underestimation
BTR	1.Noah	Niic	The largest overestimation
	2.CLM	No significant	The middle overestimation
	3.SSib	difference	The smallest overestimation
RAD	1. gap=F(3D, cosz)	Niic	Large overestimation
	2. gap=0	No significant	Slight overestimation
	3. gap=1-FVEG	difference	underestimation
TBOT	1. Zero flux	No significant	large
	2.Noah	difference	small

# 576 <u>5 Discussions</u>

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# 5.1 Major advantages of NMH-CS

The original intention of developing a Noah-MP model with the C# programming language was to analyze and edit Noah-MP code in a more efficient way, as there are many modern and efficient tools

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580 available for analyzing code written in C#, such as Microsoft Visual Studio, SharpDevelop 581 (https://github.com/icsharpcode/SharpDevelop). There are almost no comparable powerful tools for 582 analyzing FORTRAN code. This advantage is significant from the developers' perspective. 583 From the user's perspective, NMH-CS run on windows (although it should also run on other UNIX like 584 platforms after some specific configuration in the future), which is more favorable for many Windows 585 users around the world. On the windows system, in most cases, the NMH-CS software can be 586 distributed at multiple computers by simply copying it, unlike that in Unix-like systems, compiling of 587 the code is usually required.

#### 5.3 Inconsistencies between the two models

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(WRF-Hydro and NMH-CS) was found to be related to the transition between the presence or absence of snow layers. In Noah-MP, the existence of snow layers is determined by the depth of the snow (represented by the variable SNOWH in the code). When the SNOWH value approaches the threshold (0.025m), a small error will result in a division on the judgement whether a snow layer exists. This inconsistent division will further lead to significant differences in other state variables. However, this difference will not last long (as shown in Figure 3, up to 30-40 days). However, it is difficult to determine whether the errors in SNOWH are caused by the accumulation of floating-point errors or other errors.

There may be some other inconsistencies that has not been identified. Due to the complex nature of Noah-MP, it is challenging to identify all the minor differences through the process of code checking and debugging. Therefore, to ensure the results of two models to be completely consistent need a long-term process. Discrepancies can be arisen from multiple factors, including floating-point calculation errors, some inconsistent hardcoded parameter values (as local variables in certain modules), or inconsistent process are discrepancies as discrepancies and acceptable values are a sectable values are discrepancies.

As indicated by the previous analysis, the main inconsistency between the outputs of the two models

errors, some inconsistent hardcoded parameter values (as local variables in certain modules), or inconsistent programming code. The former two are reasonable and acceptable, whereas a coding mismatch can cause unexpected outputs. From the scientific perspective, these minor differences between NMH-CS and WRF Hydro are not very critical, as the model users are always modifying the code during their research, and small changes in the code can lead to large different results. The existence of differences does not always mean that NMH-CS is inferior to Noah-MP in WRF-Hydro 3.0.

In most cases, identifying discrepancies is only feasible during the debugging of the first 1-3 timesteps,

but not for tens to hundreds of subsequent iterations. It is not uncommon for errors to remain
undetected even after the execution of numerous time steps. In this study, given that no code
inconsistencies were found after multiple rounds of code checking, it is plausible that floating-point
errors related to SNOWH (or other related variables) play a major role in explaining the remaining
discrepancies.

Based on the debugging process, we also found that some variables such as TV and TG calculated by
the two models always have slight inconsistencies, but they are almost insignificant on a daily or
monthly scale. It is highly probable that such inconsistencies arise from the accumulated error caused
by the recursive calculation of energy transferring for vegetated and bare land.

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621	65 Model code and technical documentation for NMP-HydroNMH-CS
622	We archive, manage, and maintain the NMP-HydroNMH-CS at https://github.com/lsucksis/NMP-
623	HydroNMH-CS for public access. A technical description was provided at the same site. The original
624	version of the model is also provided at the website of Science Data Bank:
625	https://doi.org/10.57760/sciencedb.16102.
626	75 Conclusions
627	This study presents the NMP-HydroNMH-CS 31.0, which is a reconstructed land surface eco-
628	hydrological model based on Noah-MP. –The model was developed by translating the FORTRAN
629	code of Noah-MP (in WRF-Hydro 3.0) into C# and also coupling it with a river routing model. The
630	model has been designed for parallel execution on Windows systems, thereby capitalizing on the multi-
631	core CPUs that are now a standard feature of personal computers. The NMP-HydroNMH-CS code has
632	been subjected to rigorous testing to ensure that it produces results that are <u>as</u> _consistent <u>as possible</u>
633	with those of the original WRF-Hydro. The code is based on the C# language, which facilitates greater
634	user-friendliness and facilitates modification and expansion.
635	The development of this software enabled the successful execution of high-resolution simulations
636	encompassing a 6-km span within the Yellow River Basin (YRB). These simulations were conducted
637	with a multitude of parameter scheme (PS) combinations within the Noah-MP framework. Maps of all
638	the outputs (runoff, evaporation, groundwater, energy, vegetation) across the grid domain demonstrate
639	consistent spatial patterns that are simulated by the two models. The long-term variations of multiple
640	state variables simulated by the two models also exhibit high consistency, although some differences
641	are evidentalso exist. Identifying the cause of this simulation discrepancy in the outputs of the two-
642	models is a challenging task, given the intricate nature of the Noah-MP code. The sporadic occurrence-
643	of errors may be attributed to the accumulation of floating-point numerical calculation errors,
644	especially for the cases below frozen temperature.
645	With regard to the By enabling the coupled river routing model on river networkLanzhou hydrological
646	station, the river discharge simulated by NMP-HydroNMH-CS 13.0based on the multiple scheme
647	combination of parameterizations is found to be in reasonable agreement with the reconstructed natural
648	river discharge, for the Lanzhou hydrological station.
649	The main inconsistencies in multiple variables between NMH-CS output and WRF-Hydro output was
650	found to be related to inconsistent judgments on the presence of snow layers, which are caused by
651	minor cumulative errors near the threshold value of 0.025m for snow depth. Overall, while there are
652	occasionally happened disparities in the NMH-CS simulated results when compared to the original
653	WRF-Hydro, it reproduces highly consistent spatiotemporal distribution of multiple variables as that by
654	WRF-Hydro 3.0. It can therefore be asserted that NMH-CS can be considered a reliable replica of
655	Noah-MP in the uncoupled WRF-Hydro 3.0.
656	Overall, while there are discrepancies in the simulated results when compared to the original model, the
657	NMP-Hydro 1.0 reproduces consistent spatiotemporal distribution of multiple variables as that by WRF-
658	Hydro 3.0. Given the complex nature of long-term state variables in Noah-MP, which reflect multiple
659	processes including runoff production, energy transfer and dynamical vegetation, the results of NMP-
660	Hydro and WRF-Hydro/Noah-MP remain highly consistent. It can therefore be asserted that NMP-Hydro
661	ean be considered a reliable replica of Noah-MP in the uncoupled WRF-Hydro 3.0. It was inevitable that

minor modifications to the code or model parameters would be required during the testing phase of NMP-Hydro. This presented a significant challenge in reproducing the identical outputs as those generated by WRF-Hydro. This new software NMH-CS can run on Windows system platforms. Its C# code can be analyzed and visually browsed using many modern intelligent tools such as those in Sharpdevelop or Microsoft Visual Studio. This feature makes the code easier to analyze and modify, which in turn will attract more users and promote the future development of the Noah-MP model. The current version of NMH-CS can serve as a good model for simulating land surface processes in climate change and ecohydrology research. Although NMH-CS cannot be used as a coupling module to other FORTRAN based framework models (such as the WRF model), it can still be used as a prototype system to improve the Noah-MP schemes. Any new improvements in NMH-CS can easily be updated to other FORTRAN based Noah-MP.

Future plans for the development of NMH-CS include (1) providing a single-column run mode and incorporating a genetic algorithm-based parameter optimization module; (2) extending the functionality for modelling dynamic vegetation by designing new schemes or optimizing parameters; (3) implementing major improved model physics that exists in later versions (for example the Noah-MP 5.0) of Noah-MP into the NMH-CS framework; (4) enabling the functionality of running on UNIX-like systems.

678 <u>systems</u>

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682 intricacies of Noah-M *Code and data a* 

Code and data availability. 1. The <a href="MMP-HydroNMH-CS">MMP-HydroNMH-CS</a> model code is available at <a href="https://doi.org/10.57760/sciencedb.16102">https://doi.org/10.57760/sciencedb.16102</a> (Liu, 2024b)(Liu, 2024a). 2. The Noah-MP technical documentation is available at the same site and more details will continue to be added in the documentation. 3. The benchmark meteorological datasets for driving <a href="https://doi.org/10.57760/sciencedb.13122">NMP-HydroNMH-CS</a> and WRF-Hydro 3.0 were uploaded to the Science Data Bank (DOI: <a href="https://doi.org/10.57760/sciencedb.13122">https://doi.org/10.57760/sciencedb.13122</a> (Liu, 2024b)(Liu, 2024a)).

Author contributions. Yong-He Liu has translated the code of WRF-Hydro/Noah-MP to NMP-HydroNMH-CS, the debugging and the benchmark model simulations. The work is led by Zong-Liang Yang. Liu has drafted the paper, with improvement made by Yang.

Competing interests. The contact author has declared that none of the authors has any competing interests.

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