

PIBM 1.0: An individual-based model for simulating phytoplankton acclimation, diversity, and evolution in the ocean

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Abstract. Phytoplankton is a diverse group of photosynthetic organisms and accounts for almost half of global primary production. However, most existing marine ecosystem models incorporate limited phytoplankton diversity, overlook phytoplankton evolution, and treat phytoplankton as concentrations instead of particles. Here we present an individual-based phytoplankton model that captures three dimensions of phytoplankton traits (size, temperature, and light affinities) and allows phytoplankton cells to mutate in a one-dimensional (1D) water column. Other ~~components of this ecosystem~~ ecosystem components include dissolved inorganic nitrogen, ~~twenty-size~~ twenty-size classes of zooplankton, and detritus, all modelled as Eulerian fields. This hybrid plankton model can not only reproduce the general seasonal patterns of nutrients, chlorophyll, and primary production in the subtropical ocean, but also simulate variations in phytoplankton traits and functional diversity. We expect that this model will be a useful tool for studying phytoplankton physiology, diversity and evolution in the ocean.

10 *Copyright statement.*

1 Introduction

The ocean carbon cycle plays a pivotal role in ~~affecting how~~ shaping the Earth System ~~responds~~ 's response to climate change (Sigman and Boyle, 2000). ~~Phytoplankton is one of the most critical player in the global ocean carbon cycle, particularly,~~ with phytoplankton as one of its most critical players. Phytoplankton are essential in transporting carbon from the surface to the deep ocean, a process known as the ~~biological carbon pump~~ biological carbon pump (Ducklow et al., 2000; Sigman and Boyle, 2000). ~~In addition,~~ phytoplankton constitutes the basis biological carbon pump (Ducklow et al., 2000; Sigman and Boyle, 2000). ~~Additionally,~~ phytoplankton form the foundation of the marine food web and ~~contributes to almost~~ contribute to nearly half of global primary production (Field et al., 1998). ~~Given that numerical models are our primary tool to predict how~~ Therefore, accurate predictions of the Earth System ~~will respond to climate change induced by anthropogenic release of carbon dioxide, the phytoplankton~~ models have to capture the main aspects of phytoplankton's response to anthropogenic carbon emissions rely on numerical models that effectively capture the physiology and ecology of phytoplankton.

~~There are a number of limitations associated with~~ Despite their importance, many existing phytoplankton models ~~First,~~ the majority of plankton have notable limitations. Most of these models use the Eulerian framework, ~~while in reality~~ which

assumes that phytoplankton cells are fixed at specific locations in the water column. However, in reality, phytoplankton cells are dispersed ~~in throughout~~ the water column by ~~turbulence~~. While a combination of turbulence, advection, mixing, and other dynamic processes. Although Eulerian models are widely used and accepted, they can ~~induce~~ introduce discrepancies due to Jensen's inequality, where the average of the mean ~~may does~~ not necessarily equal the mean of the average (Baudry et al., 2018; Christensen et al., 2022).

It is still unclear whether the Eulerian model ~~Uncertainty remains about whether Eulerian models~~ or fixed-depth ~~incubations over- in situ incubations overestimate~~ or underestimate primary production ~~if we assume the Lagrangian phytoplankton model represents the ground truth~~ compared to the realistic situation in which phytoplankton cells can move in the water column. For example, Barkmann and Woods (1996) suggested that ~~the~~ primary production estimates based on ~~incubation bottles fixed at certain depths fixed-depth incubation bottles~~ could overestimate the true primary production by ~~up to as much as 40%~~, while the computation by Ross and Geider (2009) suggested %, while Ross and Geider (2009) reported only a minor difference.

The ~~above problem becomes even murkier~~ challenge intensifies when phytoplankton acclimation is ~~taken into account due to the different time scales of acclimation and considered~~, as it can operate at a different time scale than mixing (Tomkins et al., 2020). ~~Consider the case that phytoplankton acclimate to varying light conditions by changing~~ For instance, when phytoplankton adjust their intracellular carbon-to-chlorophyll (C:Chl) ratios ~~. If the water column mixing rate is faster than the acclimation rate, phytoplankton cells are essentially acclimating to the average light condition experienced throughout in response to changing light conditions, the relative speeds of acclimation and mixing become crucial. If mixing occurs faster than acclimation, cells effectively experience an averaged light environment over their life cycle, which may exacerbate potentially amplifying the effect of Jensen's inequality. By contrast, if the mixing rate is slower than the acclimation rate, phytoplankton cells constantly~~ Conversely, when acclimation is faster than mixing, cells continuously adjust their intracellular C:Chl ratios ~~which is more similar to the Eulerian scenario in which phytoplankton cells are fixed at, mirroring an Eulerian scenario where phytoplankton remain at a fixed depth~~. In addition to light acclimation, phytoplankton can ~~also~~ adjust their intracellular nutrient quota to acclimate to the external nutrient environment (Morel, 1987). Further ~~considering the stochastic nature of the movements of phytoplankton cells, it is challenging if not impossible to quickly tell the difference between the Eulerian and Lagrangian model outputs.~~

What complicates things further is that the acclimation time scales (and other rates such as ~~complexity~~ arises because ~~acclimation timescales, along with~~ photosynthesis and nutrient uptake ~~rates~~, depend on phytoplankton traits ~~such as like~~ cell size (Litchman et al., 2009; Edwards et al., 2012). ~~When comparing small against large cells~~ For instance, acclimation may ~~take place much faster for a small cell than a large cell. Unfortunately, most of the~~ occur much faster in small cells than in large cells.

Given these limitations, Lagrangian models, also known as Individual-Based Models (IBMs), present a compelling alternative to Eulerian approaches in ecological modeling. By simulating individual interactions, adaptive behavior, and responses to environmental conditions, IBMs capture dynamics that conventional models often overlook (Woods and Onken, 1982; Grimm et al., 2006; . While the first IBMs emerged in the late 1970s and early 1980s (Ledbetter, 1979; Platt and Gallegos, 1981; Falkowski and Wirick, 1981; V , their widespread use was limited due to high computational costs. Advances in computing power, however, have since made

module includes three traits, cell size, optimal temperature, and light affinity, while also incorporating phytoplankton acclimation capability and evolutionary dynamics. This type of ~~mixed-hybrid~~ Eulerian and Lagrangian ~~modelling-modeling~~ approach has been used beyond plankton ~~modelling-modeling~~, such as in the field of aerosol-cloud ~~interactions-interaction~~ (Grabowski et al., 2019; Dziekan and Zmijewski, 2022).

~~The model's primary currency is nitrogen, but it also estimates the carbon and chlorophyll content of phytoplankton cells. Phytoplankton physiological rates, which characterise each super-individual, vary with time and depth according to nutrient availability, temperature conditions, and light (Geider et al., 1998; Ross and Geider, 2009), and three master traits: cell size (expressed in terms of the maximal carbon content per cell during its life cycle), optimal temperature, and light affinity (expressed in terms of the initial slope of the Photosynthesis-Irradiance (P-I) curve). The diversity of phytoplankton super-individuals is sustained by grazing and mutation. The model structure is shown in Fig. 1.~~

In the following sections, we first describe our ecological model and the differential equations that govern the growth and selection of phytoplankton. Next, we present the main results of the model and discuss its merits and limitations. To validate the model outputs, we compare them with *in situ* observations at the Bermuda Atlantic Time-series Study (BATS) station. Moreover, we compare PIBM outputs with those of two versions, Eulerian and Lagrangian, of a simple Nitrogen-Phytoplankton-Zooplankton (NPZD) model. This allows us to examine the effects of increased model complexity to understand how incorporating multiple phytoplankton traits influences biomass, production, and the differences between Eulerian and Lagrangian approaches. It is important to note that the objective of developing PIBM and performing model validation/comparison is not to improve simulations of bulk properties (e.g., nutrient and chlorophyll concentrations) against simpler models but to provide a tool to better study phytoplankton diversity in the ocean.

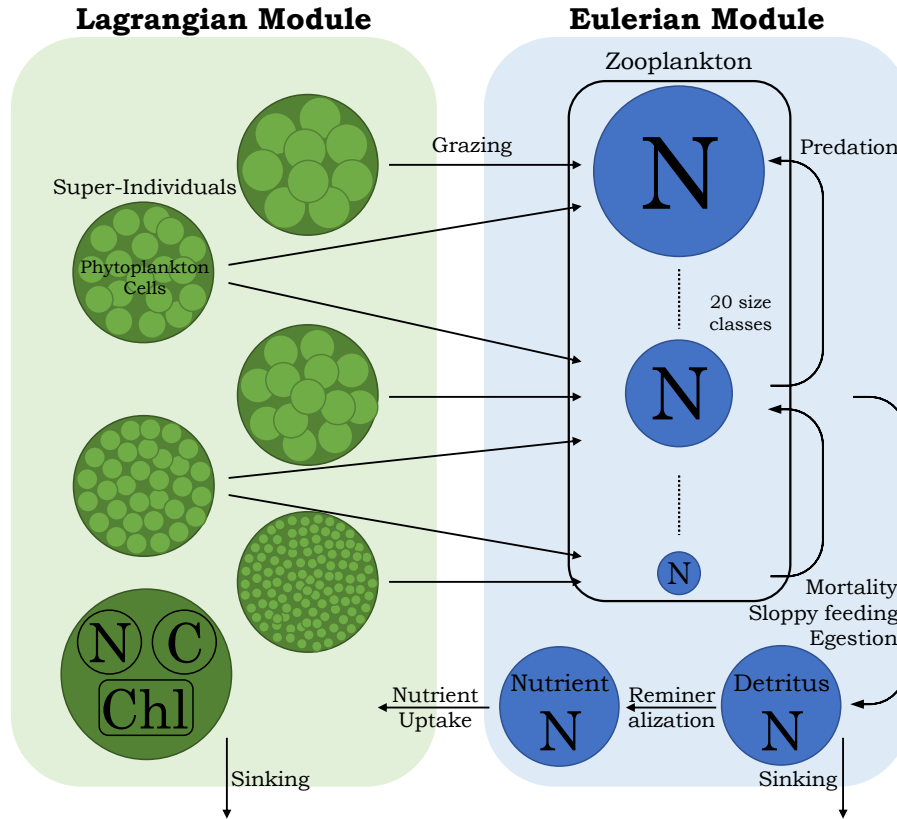


Figure 1. Conceptual diagram of the ~~1D-hybrid-Eulerian-Lagrangian-1D PIBM~~ model. N: Nitrogen. C: Carbon. Chl: Chlorophyll. The black arrows represent nitrogen flows.

2 Model description

2.1 Overview

115 ~~The model PIBM is a 1D hybrid model expanded on the classic Nitrogen-Phytoplankton-Zooplankton-Detritus (NPZD) model. Phytoplankton-Eulerian-Lagrangian hybrid system that extends the classic NPZD framework. In PIBM, phytoplankton cells are represented by super-individuals (the Lagrangian module) see below) within the Lagrangian module, which is coupled to an Eulerian module, that Eulerian module. The Eulerian module calculates the dynamics of dissolved inorganic nitrogen (DIN), multiple size classes of zooplankton (ZOO), and detritus (DET) as continuous concentrations along across the vertical domain.~~

120 ~~In the next subsections, we will describe both modules in detail. All model parameters are While nitrogen is the model's primary currency, the model also estimates the carbon and chlorophyll content of phytoplankton cells. The model structure is illustrated in Fig. 1, with all model parameters listed in Tables 1 and 2. The following subsections provide detailed descriptions of both modules.~~

2.2 Lagrangian module

2.2.1 Super-Individuals

The *Lagrangian* module simulates the ~~phytoplanktonic community as~~ phytoplankton community using super-individuals ~~(a cluster, which represent clusters of identical phytoplankton cells) to represent~~. This approach allows us to model a realistic number of ~~phytoplankton cells with affordable computational costs~~ cells while keeping computational costs manageable (Scheffer et al., 1995). To avoid memory ~~issues, we assign a constant~~ limitations, the number of super-individuals remains
130 constant throughout the simulation.

2.2.2 Phytoplankton model

~~We assign~~

Phytoplankton physiological rates, such as growth and nutrient uptake, are determined by three master traits: cell size, optimal temperature (T_{opt} , °C), and the initial slope of the Photosynthesis-Irradiance (P-I) curve (α^{Chl} , (W² m⁻¹ g Chl (mol C)⁻¹)⁻¹ d⁻¹) to each phytoplankton super-individual. Cell size determines the capability and light affinity. These rates vary over time and depth, responding not only to changes in the external environment (nutrient availability, temperature, and light) but also to changes in community composition (traits) (Geider et al., 1998; Ross and Geider, 2009).
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Cell size influences the ability of phytoplankton to take up inorganic nutrients and its vulnerability to zooplankton grazing. Cell size is expressed in terms of It is expressed as the maximal carbon content per cell (C_{div} , pmol C cell⁻¹). Note that while the, which reflects the phytoplankton size in terms of its maximum carbon content during the division phase. While actual cellular carbon content (P_C , pmol C cell⁻¹) can vary as a result of fluctuates due to photosynthesis and respiration depending on its nutrient and light environment, we assume that the traits of, nutrient uptake and zooplankton grazing only depend grazing rates depend solely on C_{div} because otherwise these traits will change constantly with time. This simplification ensures that these rates are not constantly recalculated due to changes in P_C and only change when C_{div} varies due to mutation. Optimal temperature (T_{opt} , °C) and light affinity, expressed as the initial slope of the photosynthesis-irradiance curve (α^{Chl} , (W² m⁻¹ g Chl (mol C)⁻¹)⁻¹ d⁻¹), describe phytoplankton response to environmental temperature and light conditions.
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We assign the three master traits to each phytoplankton super-individual at the beginning of the simulation, with values randomly drawn from a uniform distribution within predefined minimum and maximum limits. Once assigned, trait values remain constant throughout the simulation unless a mutation occurs (Section 2.2.3), which helps to sustain diversity.

~~The dynamics of~~ At the heart of the Lagrangian module lie the equations governing phytoplankton cellular carbon (P_C), nitrogen (P_N , pmol N cell⁻¹), and chlorophyll (P_{Chl} , pg Chl cell⁻¹) are modelled following Geider et al. (1998):, formulated following Geider et al. (1998). Geider et al. (1998)'s equations describe the balance of photosynthesis, nutrient uptake, respiration, and chlorophyll synthesis, capturing how phytoplankton regulate their cellular contents in response to environmental conditions. Eq.1 states that phytoplankton cellular carbon is fueled by photosynthesis, which converts inorganic carbon into organic carbon, but is depleted by the cost of nutrient uptake and respiration. Eq.2 states that phytoplankton cellular nitrogen is fueled
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by nitrogen uptake but is depleted by respiration. Eq. 3 states that phytoplankton cellular chlorophyll content is fueled by chlorophyll synthesis, which depends on both photosynthesis and nitrogen uptake, but is consumed by respiration.

$$\frac{1}{P_C} \frac{dP_C}{dt} = P^C - \zeta V_N^C - R_C f(T), \quad (1)$$

$$160 \quad \frac{1}{P_N} \frac{dP_N}{dt} = \frac{V_N^C}{Q^N} - R_N f(T), \quad (2)$$

$$\frac{1}{P_{Chl}} \frac{dP_{Chl}}{dt} = \frac{\rho_{Chl} V_N^C}{\theta^C} - R_{Chl} f(T) \quad (3)$$

P^C (d^{-1}) is the carbon-specific photosynthesis rate, ζ ($mol\ C\ mol\ N^{-1}$) is the cost of biosynthesis, V_N^C ($mol\ N\ mol\ C^{-1}\ d^{-1}$) is the nitrogen uptake rate, Q^N ($mol\ N\ mol\ C^{-1}$) is the cellular N:C ratio, ρ_{Chl} (dimensionless) is the fraction of phytoplankton carbon production that is devoted to chlorophyll synthesis, θ^C ($g\ Chl\ mol\ C^{-1}$) is the ratio of Chl synthesis to carbon fixation (representing phytoplankton acclimation to light variability), R_C , R_N and R_{Chl} (d^{-1}) are the phytoplankton respiration rates for carbon, nitrogen and chlorophyll, respectively. $f(T)$ is the function describing the temperature (T , K) dependence of phytoplankton metabolism which is detailed in the temperature section below.

~~Eq. 1 states that phytoplankton cellular carbon is fueled by photosynthesis that converts inorganic carbon to organic carbon, but is depleted by the cost of both nutrient uptake and respiration. Eq. 2 states that phytoplankton cellular nitrogen is fueled by phytoplankton uptake of nitrogen, but is depleted by respiration. Eq. 3 states that phytoplankton cellular chlorophyll content is fueled by chlorophyll synthesis which depends on both photosynthesis and nitrogen uptake, but is consumed by respiration.~~

It is important to note that we implicitly assume that P_C , P_N , and P_{Chl} are not affected by zooplankton grazing which only reduces the number of cells per super-individual.

175 P^C is a function of light availability as defined below:

$$P^C = P_m^C \left[1 - e^{\left(\frac{-\alpha^{Chl} \theta^C I A^\infty}{P_m^C \text{max}} \right)} \right] \quad (4)$$

where P_m^C (d^{-1}) is the maximal carbon-specific photosynthesis rate, I ($W\ m^{-2}$) is the irradiance, and A^∞ (dimensionless) is the term that accounts for photosynthetic photoinhibition (*see* Eq. 6) ~~which was not~~. Although photoinhibition was not originally included in Geider et al. (1998), incorporating it is essential for capturing the decline in photosynthetic efficiency under overly strong irradiance for species that are adapted to low light levels (Moore et al., 1998; Han, 2002; Ross et al., 2011)

P_m^C depends on intracellular nutrient status:

$$P_m^C = \mu_m \frac{Q_{max}^N - Q_{min}^N}{Q_{max}^N - Q_{min}^N} \quad (5)$$

where μ_m (d^{-1}) is the maximal specific growth rate ~~as a function of temperature under resource (nutrient and light) replete~~
 185 ~~conditions (see , which varies with temperature under nutrient- and light-replete conditions (Eq. 10), and ,~~ Q_{max}^N (mol N mol
 C^{-1}) is the maximal nitrogen-to-carbon ratio and Q_{min}^N is the minimal nitrogen cell quota. ~~Q_{max}^N and Q_{min}^N are size dependent~~
~~; both depend on cell size (Table 2). Here, we replaced the original P_{ref}^C (maximal photosynthetic capacity at a reference~~
~~temperature) from Geider et al. (1998) with μ_m which depend on both temperature and light traits as shown later.~~

To simulate the short-term responses of phytoplankton cells to potential high light stress when ~~being they are~~ dispersed
 190 to the surface, we include photoinhibition into the phytoplankton model following ~~Ross and Geider (2009)~~ ~~previous studies~~
~~(Han, 2002; Ross et al., 2011; Nikolaou et al., 2016)~~. Photoinhibition decreases the photosynthesis rate due to ~~the damage of~~
 D1 protein ~~damage~~ under high light, and it is expressed as the fraction of open Photosynthetic Units (PSU) (Han, 2002;
 Nikolaou et al., 2016):

$$A^\infty = \frac{1}{1 + \sigma_{PSII} I \tau + K \sigma_{PSII}^2 I^2 \tau} \quad (6)$$

195 σ_{PSII} ($m^2 W^{-1}$) is the effective absorption cross-section of ~~the~~ Photosystem II (PSII) ~~and is ,~~ parameterized as a power-law
 relationship with θ^C : $\sigma_{PSII} = \delta (\theta^C)^\kappa$ (Nikolaou et al., 2016). ~~Here, δ ($(W m^{-2})^{-1} (g Chl g C^{-1})^{-1}$) and κ (dimensionless)~~
 are constants (Table 1). τ (s) is the turnover time of the electron transfer chain and K (s^{-1}) is the ratio of damage to repair
 constants ($K = k_d/k_r$).

The tension between ~~photo-damage-photodamage~~ (k_d , dimensionless) and repair (k_r , s) of a PSU determines the fraction of
 200 open reaction centres and ~~abundances the abundance~~ of D1 proteins at a given light level. To set up a trade-off between high-
 light adapted and low-light adapted ecotypes (Moore et al., 1998), we assume that k_r and α^{Chl} are negatively correlated ~~such~~
~~that phytoplankton cells that are , Thus, phytoplankton cells~~ adapted to low-light (i.e., with larger α^{Chl}), have a reduced
 capability of photo-repair (i.e., smaller K) (Key et al., 2010). Conversely, phytoplankton cells adapted to high-light (larger
 K) ~~have a greater capability are better able~~ to cope with photoinhibition but are less efficient in absorbing photons under low
 205 light (smaller α^{Chl}). In addition, we consider the effect of nutrient status ~~of phytoplankton on cells' ability to photo-repair on~~
~~their ability to perform photorepair~~, as nutrient limitation can jeopardize photosynthetic energy transfer efficiency (Herrig and
 Falkowski, 1989). ~~Therefore, we come up with~~ ~~Based on this, we propose~~ the following relationship ~~describing for~~ k_r :

$$k_r = a \left(\frac{\alpha^{Chl}}{b} \right)^v \frac{Q_{max}^N - Q_{min}^N}{Q_{max}^N - Q_{min}^N}, \quad (7)$$

~~in which where~~ $a = 2 \times 10^{-5}$, $b = 5 \times 10^{-7}$, and $v = -6.64$ are constants. These parameters ~~are obtained via fitting to the data~~
 210 ~~of were obtained by fitting data from~~ *Prochlorococcus* in Moore et al. (1998).

Phytoplankton nitrogen uptake (V_N^C) depends on both external ~~levels of DIN~~ dissolved inorganic nitrogen (DIN, mmol N m^{-3}) levels and intracellular nitrogen status:

$$V_N^C = V_m^C \frac{DIN}{DIN + K_N} \left(\frac{Q_{max}^N - Q^N}{Q_{max}^N - Q_{min}^N} \right)^n \quad (8)$$

with where $V_m^C = \mu_m Q_{max}^N$ (mol N mol C⁻¹ d⁻¹) being is the maximal specific nitrogen uptake rate, and K_N (μM) being is the half-saturation constant for DIN uptake. The parameter n (dimensionless), varying which varies between 0 and 1, is the shape factor adjusting determines the dependence of the maximum uptake rate (V_m^C) on cell quota (Q^N) (Geider et al., 1998). All three parameters K_N , Q_{max}^N , and Q_{min}^N depend on cell size following allometric relationships (Table 2).

ρ_{Chl} (dimensionless) depends on light, photosynthetic rate, the initial slope of the P-I photosynthesis-irradiance curve (α^{Chl}), and Chl:C ratio (θ^C):

$$\rho_{Chl} = \theta_{max}^N \frac{P^C}{\alpha^{Chl} \theta^C I} \quad (9)$$

where θ_{max}^N is the maximum chlorophyll to nitrogen chlorophyll-to-nitrogen ratio (g Chl mol N⁻¹). During dark hours, when $I = 0$, ρ_{Chl} is assumed to be equal to retain the value calculated at the end of the preceding light period.

The maximal growth rate, μ_m , depends on T_{opt} (K) as well as the environmental temperature (T , K) following (Chen, 2022) which extends from, following Chen (2022), which extends the Metabolic Theory of Ecology (Dell et al., 2011; Chen and Laws, 2017):

$$\mu_m = \mu_0 \frac{E_a + E_d}{E_d} \frac{e^{E_a(x-\theta)}}{1 + \frac{E_a}{E_d} e^{(E_a + E_d)(x-\theta)}} \quad (10)$$

in which where θ and x are the transformed optimal and environmental temperatures (for mathematical convenience), respectively ($\theta = \frac{1}{k_b} \left(\frac{1}{T_0} - \frac{1}{T_{opt}} \right)$ and $x = \frac{1}{k_b} \left(\frac{1}{T_0} - \frac{1}{T} \right)$);

$$\theta = \frac{1}{k_b} \left(\frac{1}{T_0} - \frac{1}{T_{opt}} \right), \quad x = \frac{1}{k_b} \left(\frac{1}{T_0} - \frac{1}{T} \right) \quad (11)$$

with T_0 (K) being the reference temperature set at 288 K. μ_0 (d⁻¹) is the normalization constant for μ_m (at $\theta = 0$), E_a (eV) is the intraspecific activation energy, and E_d (eV) is the nominal activation energy regulating how fast the growth rate (μ , d⁻¹) decreases with increasing T when $x > \theta$. Finally, k_b (eV K⁻¹) is the Boltzmann constant.

Based on analyzing a dataset of phytoplankton growth against the analysis of a dataset relating phytoplankton growth to temperature, μ_0 (d⁻¹), E_a (eV), and E_h (eV) are found to be allometric functions of θ (Chen, 2022):

$$\mu_0 = \mu' e^{E_i \theta} \quad (12a)$$

$$E_a = E_{a0} e^{\beta \theta} \quad (12b)$$

$$E_h = E_{h0} e^{\phi \theta} \quad (12c)$$

in which μ' (d^{-1}) is the normalization constant for μ_0 when $\theta = 0$, E_i (eV) is the interspecific activation energy, E_{a0} (eV) is the normalization constant for E_a when $\theta = 0$, β (eV) is the scaling exponent against θ for E_a , E_{h0} (eV) is the normalization constant for E_h when $\theta = 0$, and ϕ (eV) is the scaling exponent for E_h .

Following the mechanistic model of Wirtz (2011), we ~~assume that let~~ μ' ~~varies to decrease~~ with cell Equivalent Spherical Diameter (ESD, μm) as a result of intra-cellular self-shading and excess density:

$$\mu' = \frac{\mu'_0}{1 + a_0 \left(\frac{\rho^s}{\rho_0} \right)^{\frac{1}{3}} \text{ESD}} \quad (13)$$

where a_0 ($= 0.34 \text{ m}^{-1}$) is the length scale of photosynthesis depletion, ρ^s ($= 0.25 \text{ pg C m}^{-3}$) is the carbon density at a reference ESD ($ESD_s = 8.00 \mu\text{m}$), and ρ_0 ($= 0.50 \text{ pg C m}^{-3}$) is the specific carbon density for the relative chloroplast volume (V_{Chl}/V).

We also ~~assume consider~~ that phytoplankton respiration (R_C , R_N , and R_{Chl}) ~~follow follows~~ the same temperature dependence as μ_m (Barton et al., 2020).

Following Wirtz (2011), ~~we assume that~~ these specific respiration rates are assumed to scale inversely with phytoplankton ESD:

$$R_C = R_{C,s} \frac{\text{ESD}_s}{\text{ESD}} \quad (14)$$

where $R_{C,s}$ is the ~~temperature-dependent temperature-dependent~~ respiration rate at ESD_s . If $T = T_{opt}$, $R_{C,s} = 0.025 \text{ d}^{-1}$ for the specific respiration rates of phytoplankton carbon, nitrogen and chlorophyll.

Eqs. 13 and 14 capture the unimodal relationship between maximal growth rate and cell size (Chen and Liu, 2010, 2011; Marañón et al., 2013). If a phytoplankton cell is too large, its self-shading and intracellular decline of CO_2 reduces its maximal growth rate. On the other hand, if a cell is too small, the high specific respiration cost leads to a rapid decline in maximal growth rate. This ~~constraint tradeoff~~ on the range of cell size plays a ~~fundamental role in shaping phytoplankton size structure~~ critical role in constraining phytoplankton size range in the model in which the phytoplankton cells are allowed to evolve freely by mutation. Without this tradeoff, phytoplankton will evolve towards infinitely small cells in very oligotrophic environments.

2.2.3 Phytoplankton division, ~~death, mutation~~ and ~~evolution death~~

Phytoplankton cell division ~~takes place occurs~~ when the cellular carbon content reaches C_{div} (Cianelli et al., 2009; Ross and Geider, 2009). ~~When a cell divides~~ Upon division, the parent cell is split into two equal daughter cells, each inheriting half of the carbon, nitrogen, and chlorophyll content. As a consequence, the number of cells per super-individual doubles.

~~Phytoplankton cell dies when phytoplankton cellular carbon falls below a minimal threshold (C_{min} , pmol C cell^{-1}) defined as a quarter of C_{div} (Ross and Geider, 2009) or when the total amount of nitrogen of a super-individual drops below 0.10% of the average nitrogen content of all super-individuals.~~

~~It is important to note that in addition to respiratory cost, phytoplankton cells suffer from zooplankton grazing. We assume that zooplankton grazing can only reduce the number of cells per super-individual without affecting cellular carbon or nitrogen.~~

When a cell dies, its nitrogen content is converted into detritus. At the same time, to keep the number of super-individuals constant, By default, the super-individual with the maximum nitrogen content is divided into two new super-individuals each with half of the number of cells of the parent super-individual.

When phytoplankton cells divide, they are allowed to daughter cells inherit the same trait values for C_{div} , T_{opt} , and α^{Chl} as the parent cell; however, they have a small probability to mutate (i.e., changing the values of the three traits $\log(C_{div})$, T_{opt} , and $\log(\alpha^{Chl})$) following of mutation, with the trait values altered according to a Gaussian distribution with the mean equal to the parent cell's trait's value and a given specified standard deviation. The probability of mutation for a super-individual to mutate ($\nu_{j,phy}$) is proportional to its number of cells. Although it ($n_{j,phy}$), such that $\nu_{j,phy} = \nu_0 n_{j,phy}$, where ν_0 represents the mutation rate per cell.

Although this mutation framework may not accurately reflect reality, for the sake of simplicity and ease of modelling, we assume that all ecotypes share the same mutation rate. However, in reality, mutation rates can vary among different phytoplankton ecotypes (Beardmore et al., 2011), and even within the same species when subjected to stress (Bjedov et al., 2003). Besides Additionally, while we assume that the mutation of one trait is independent of others, the user can modify the mutation covariance matrix to change how the mutation of one trait can depend depends on those of other traits.

2.2.4 Mean trait and trait (co)variance

We characterize phytoplankton community composition in terms of mean trait and trait covariance. Trait covariance can be used to represent part of functional diversity (Norberg et al., 2001; Chen et al., 2019).

The community mean phytoplankton trait is calculated as the carbon-weighted mean of all phytoplankton cells in the community:-

$$\bar{l} = \frac{\sum_{i=1}^k l_i n_i P_{C,i}}{\sum_{i=1}^k n_i P_{C,i}}$$

in which l_i and n_i represent the trait value and the number of cells of super-individual i , respectively. $P_{C,i}$ represents the cellular carbon content of Phytoplankton cells die when their cellular carbon content falls below a minimal threshold (C_{min} , $\mu\text{mol C cell}^{-1}$), defined as a quarter of C_{div} (Ross and Geider, 2009), or when the total nitrogen content of a super-individual i .

The trait covariance is calculated as:-

$$\text{COV}(l_j, l_m) = \frac{\sum_{i=1}^k (l_j - \bar{l}_j) (l_m - \bar{l}_m) n_i P_{C,i}}{\sum_{i=1}^k n_i P_{C,i}}$$

where \bar{l}_j and \bar{l}_m represent the mean trait value of trait j and m , respectively drops below 0.10% of the average nitrogen content of all super-individuals. It is important to note that we treat $\log(C_{div})$, T_{opt} , and $\log(\alpha^{Chl})$ as traits as they are more likely to follow normal distribution than the raw units.

2.2.4 Functional diversity

We calculated the functional richness and evenness of phytoplankton cells using the R-package **TPD** which considers intraspecific trait variability by computing multidimensional trait probability densities using Gaussian kernels (Mason et al., 2005; Carmona et al., 2019).

In brief, the functional richness is calculated as the amount of functional space occupied by all phytoplankton superindividuals in a community, similar to the volume calculated by the hypervolume approach (Blonder et al., 2018). The functional evenness measures how even the relative abundances of different trait values are in a community. In addition to respiratory cost, phytoplankton cells are subject to zooplankton grazing. We assume that zooplankton grazing can only reduce the number of cells per super-individual without affecting cellular carbon or nitrogen. When a cell dies, its nitrogen content is converted into detritus. Simultaneously, to maintain a constant number of super-individuals, the super-individual with the maximum nitrogen content is divided into two new super-individuals, each containing half of the cells of the parent super-individual.

2.3 Eulerian module

Dynamics of dissolved inorganic nitrogen, zooplankton, and detritus are modelled as Eulerian fields.

2.3.1 Dissolved inorganic nitrogen

The temporal and spatial variability of dissolved inorganic nitrogen (DIN, mmol N m^{-3}) depends on is governed by several processes: the total nutrient uptake by phytoplankton (P_{uptake} , mmol N m^{-3} ; see Eq. 16), zooplankton excretion (Z_{exc} ; see Eq. 17) and, detritus regeneration (D_{reg} ; see Eq. 18), and vertical diffusion (last term of the last term in Eq. 15):

$$\frac{\partial \text{DIN}}{\partial t} = -P_{\text{uptake}} + Z_{\text{exc}} + D_{\text{reg}} + \frac{\partial}{\partial z} \left[K_v(z) \frac{\partial \text{DIN}}{\partial z} \right] \quad (15)$$

where $K_v(z)$ ($\text{m}^2 \text{s}^{-1}$) is denotes the vertical eddy diffusivity at each depth layer. P_{uptake} at in grid s during over the time step Δt is defined as the sum of the nutrients taken up by all super-individuals within grid s that grid:

$$P_{\text{uptake}} = \frac{1}{H(s)} \sum_{i=1}^k \left[\left(P_{N,i}(t + \Delta t) - P_{N,i}(t) \right) n_i(t + \Delta t) \right] \quad (16)$$

in which $H(s)$ (m) is the height of the vertical grid s , $P_{N,i}(t + \Delta t)$ and $P_{N,i}(t)$ (mmol N per cell) represent the cellular nitrogen content of super-individual i at time $t + \Delta t$ and t , respectively, and $n_i(t + \Delta t)$ represents the number of cells associated with the super-individual i at time $t + \Delta t$.

We assume that Z_{exc} is a constant fraction of the total amount of food ingested by all zooplankton:

$$Z_{\text{exc}} = (1 - \xi - \eta) \sum_{j=1}^{N_z} I_j \text{ZOO}_j \quad (17)$$

where ξ (dimensionless) is represents the gross growth efficiency of zooplankton, assumed and is assumed to be constant for each size class. η (dimensionless) is denotes the fraction of unassimilated food by zooplankton, also assumed constant for each size class. I_j (d^{-1}) is the per capita total ingestion rate of zooplankton in size class j and is elaborated, as detailed in the following section.

D_{reg} is a linear function of both detritus concentration (DET, mmol N m^{-3}) and temperature (T , $^{\circ}\text{C}$) ~~as well as temperature~~ :

$$D_{reg} = R_{dn} \text{ DET } f_h(T) \quad (18)$$

where R_{dn} (d^{-1}) is the conversion rate ~~from detritus to~~ of detritus into dissolved inorganic nitrogen, and $f_h(T)$ describes the temperature dependence of heterotrophic activities, including zooplankton grazing and detritus regeneration, formulated according to the Arrhenius equation:

$$f_h(T) = e^{\frac{E_z}{k_b} \left(\frac{1}{T_{ref}} - \frac{1}{T} \right)} = e^{\frac{E_z}{k_b} \left(\frac{1}{T_0} - \frac{1}{T} \right)} = e^{E_z x} \quad (19)$$

where E_z (eV) is the activation energy for heterotrophic processes (see Table 1).

2.3.2 Zooplankton

The code for modeling zooplankton biomass was developed drawing inspiration from the equations described by (Ward et al., 2012). The model resolves 20 ~~size classes of zooplankton spaced uniformly in log space~~ zooplankton size classes, spaced logarithmically from 0.80 to 3600 μm ESD. ~~We define the~~ The smallest size class ~~as is set at~~ 0.80 μm to ~~mimic~~ represent the smallest heterotrophic eukaryotes in the ocean ~~that~~, which predominantly feed on bacteria. The upper limit of 3600 μm is ~~selected as a~~ result of a tradeoff between providing appropriate grazers that can feed on ~~chosen as a tradeoff between ensuring appropriate grazers for~~ large phytoplankton and ~~computing demand as we wish to managing computational cost, as we~~ fix the number of zooplankton size classes as 20. Because phytoplankton cells are ~~allowed to freely evolve in free to evolve in size during~~ the simulation, ~~it is possible that some phytoplankton cells can evolve into an extremely~~ some may become so large (or small) size ~~that no zooplankton can feed on~~. ~~But we cannot design a size range of zooplankton too large because otherwise there would be too few zooplankton them. However, expanding the zooplankton size range too much would reduce the resolution~~ size classes within the realistic zooplankton size range.

The nitrogen biomass of each zooplankton size class (ZOO_j , mmol N m^{-3}) increases ~~with the amount of prey (including all as they consume prey, including~~ phytoplankton super-individuals and ~~zooplankton smaller than themselves) they can eat, but~~ is reduced by the predation from larger zooplankton in addition to the smaller zooplankton, but decreases due to predation by larger zooplankton and natural mortality.

$$\frac{\partial \text{ZOO}_j}{\partial t} = \text{ZOO}_j \xi_j \sum_{j_{prey}=1}^J I_{j,j_{prey}} - \sum_{j_{pred}=1+J}^{N_z} \text{ZOO}_{j_{pred}} I_{j_{pred},j} - \text{ZOO}_j m_z f_h(T) + \frac{\partial}{\partial z} \left(K_v(z) \frac{\partial \text{ZOO}_j}{\partial z} \right) \quad (20)$$

where m_z (d^{-1}) is the linear zooplankton mortality rate, and J is the total number of prey items ~~including all, including~~ phytoplankton super-individuals ~~within the grid and other smaller zooplankton and smaller zooplankton within the grid.~~

Zooplankton *per capita* ingestion rate ($I_{j_{pred},j_{prey}}$, d^{-1}) is calculated ~~following using~~ a sigmoidal functional response ~~depending that depends~~ on total prey biomass ($B_{j_{prey}}$, mmol N m^{-3}) (Ward et al., 2012):

$$I_{j_{pred},j_{prey}} = f_h(T) I_{j_{pred}}^{max} \frac{\phi_{j_{pred},j_{prey}} B_{j_{prey}}}{F_{j_{pred}} + K_{P,j_{pred}}} (1 - e^{-\Lambda F_{j_{pred}}}) \quad (21)$$

where $I_{j_{pred}}^{max}$ (d^{-1}) is the size-dependent maximum ingestion rate (Table 2) (Hansen et al., 1997; Ward et al., 2012). $\phi_{j_{pred}, j_{prey}}$ (dimensionless) ~~is~~ represents the palatability of prey j_{prey} for predator j_{pred} . $F_{j_{pred}}$ ($mmol\ N\ m^{-3}$) is the total prey availability for predator j_{pred} , and $K_{P, j_{pred}}$ ($mmol\ N\ m^{-3}$) is the grazing half-saturation constant of predator j_{pred} . The term $(1 - e^{\Lambda F})$ ~~represents~~ accounts for the effect of prey refuge ~~which reduces the grazing effort as~~, which reduces grazing effort when available prey becomes scarce (Ivlev, 1955; Mayzaud and Poulet, 1978). The total ingestion rate of zooplankton size class j is therefore given by $I_j = \sum_{j_{prey}=1}^J I_{j, j_{prey}}$.

The food availability for ~~the~~ zooplankton size class j_{zoo} ($F_{j_{zoo}}$) includes both phyto- and zooplankton prey and is computed as:

$$F_{j_{zoo}} = \sum_{i_{phy}=1}^k \phi_{j_{zoo}, i_{phy}} B_{i_{phy}} + \sum_{i_{zoo}=1}^{j_{zoo}-1} \phi_{j_{zoo}, i_{zoo}} B_{i_{zoo}} \quad (22)$$

~~in which where~~ k is the number of super-individuals within the vertical grid. $B_{i_{phy}} = n_{i_{phy}} P_{N, i_{phy}} / H$ and $B_{i_{zoo}}$ ($mmol\ N\ m^{-3}$) ~~are~~ represent the nitrogen biomass of the i_{phy}^{th} phytoplankton super-individual and the i_{zoo}^{th} zooplankton size class, respectively. ~~Note that~~
 Since there is no zooplankton prey for the smallest zooplankton size class ($j_{zoo} = 1$), ~~We~~, we assume that zooplankton do not feed on other zooplankton larger than ~~their own size, themselves~~ but can feed on any phytoplankton prey, ~~although feeding on a phytoplankton prey~~. ~~However, feeding on phytoplankton~~ larger than their optimal prey size is penalized ~~by the~~ due to low prey palatability, $\phi_{j_{pred}, j_{prey}}$ (dimensionless):

$$\phi_{j_{pred}, j_{prey}} = \exp \left[- \left(\ln \left(\frac{\vartheta_{j_{pred}, j_{prey}}}{\vartheta_{opt}} \right) \right)^2 \left(2\sigma_{j_{pred}}^2 \right)^{-1} \right] \quad (23)$$

where $\vartheta_{j_{pred}, j_{prey}}$ (dimensionless) is the predator:prey volume ratio, ϑ_{opt} (dimensionless) is the optimal predator:prey volume ratio (Kjørboe, 2009), and $\sigma_{j_{pred}}$ (dimensionless) is the standard deviation of the feeding kernel.

ϑ_{opt} is estimated ~~from based on~~ the optimal prey size ~~defined as ESD~~, which is defined by the relationship between ESD of the optimal prey ($ESD_{prey_{opt}}^{pred}$) ~~vs. and the~~ predator ESD (ESD_{pred}) (Hansen et al., 1994; Banas, 2011):

$$ESD_{prey_{opt}}^{pred} = 0.65 ESD_{pred}^{0.56} \quad (24)$$

As we assume that zooplankton grazing only affects the number of cells per super-individual ($n_{i_{phy}}$) ~~instead of rather than~~ cellular carbon or nitrogen, ~~it turns out that it is not a trivial task to estimate the estimating~~ changes in $n_{i_{phy}}$ ~~to conserve total nitrogen while conserving total nitrogen is not trivial~~. We assume that ~~the~~ phytoplankton mortality due to zooplankton grazing (g_{phy} , d^{-1}), ~~or i.e.~~, the proportional loss of nitrogen content, during the time step Δt , is equal for all phytoplankton cells within the same super-individual ~~(and the same and~~ for all phytoplankton cells within the same vertical layer). ~~While the deaths of phytoplankton cells can be~~. While phytoplankton cell deaths can be modeled as a binomial process, ~~we can the law of large numbers allows us to~~ assume that the loss of cell numbers within a super-individual is proportional to the grazing rate g_{phy} ~~thanks to the law of large numbers~~ (Beckmann et al., 2019). Therefore, we have:

$$n_{i_{phy}}(t + \Delta t) = n_{i_{phy}}(t) (1 - g_{phy} \Delta t) \quad (25)$$

The total phytoplankton biomass loss due to zooplankton grazing within each vertical layer (P_G , mmol N m⁻³) ~~can be~~
 390 ~~calculated as~~ is computed as follows (note that we drop the subscript of z for convenience):

$$\begin{aligned}
 P_G &= -\frac{1}{H} \sum_{i_{\text{phy}}=1}^k P_{N,i_{\text{phy}}}(t) \left(n_{i_{\text{phy}}}(t) - n_{i_{\text{phy}}}(t + \Delta t) \right) \\
 &= -\frac{1}{H} \sum_{i_{\text{phy}}=1}^k P_{N,i_{\text{phy}}}(t) \left(n_{i_{\text{phy}}}(t) - n_{i_{\text{phy}}}(t)(1 - g_{\text{phy}} \Delta t) \right) \\
 &= -\frac{1}{H} \sum_{i_{\text{phy}}=1}^k P_{N,i_{\text{phy}}}(t) n_{i_{\text{phy}}}(t) g_{\text{phy}} \Delta t \\
 &= g_{\text{phy}} \Delta t \frac{\sum_{i_{\text{phy}}=1}^k P_{N,i_{\text{phy}}}(t) n_{i_{\text{phy}}}(t)}{H} \\
 &= \sum_{j_{\text{zoo}}=1}^{N_Z} I_{j_{\text{zoo}},i_{\text{phy}}} \Delta t
 \end{aligned} \tag{26}$$

where k is the number of phytoplankton super-individuals within the vertical layer.

Since $\frac{\sum_{i_{\text{phy}}=1}^k P_{N,i_{\text{phy}}}(t) n_{i_{\text{phy}}}(t)}{H}$ ~~is~~ represents the total phytoplankton nitrogen concentration within each vertical layer (P_N),
 we can derive:

$$395 \quad g_{\text{phy}} = \frac{\sum_{j_{\text{zoo}}=1}^{N_Z} I_{j_{\text{zoo}},i_{\text{phy}}}}{P_N} \tag{27}$$

Hence, the number of cells within a super-individual at the next time step ($t + \Delta t$) ~~can be computed as~~ is given by:

$$n_{i_{\text{phy}}}(t + \Delta t) = n_{i_{\text{phy}}}(t) \left(1 - \frac{\sum_{j_{\text{zoo}}=1}^{N_Z} I_{j_{\text{zoo}},i_{\text{phy}}}}{P_N} \Delta t \right) \tag{28}$$

It is therefore important that the grazing effect ($n_{i_{\text{phy}}}(t + \Delta t)$) has to be computed before phytoplankton nutrient uptake
 (~~P_T~~ P_{uptake} , Eq. 16).

400 2.3.3 Detritus

Changes in the concentration of detritus (DET, mmol N m⁻³) are computed as:

$$\frac{\partial \text{DET}}{\partial t} = \sum_{j=1}^{N_z} \left(\eta I_j + m_z f_h(T) \right) \text{ZOO}_j - R_{dn} \text{DET} f_h(T) - W_d \frac{\partial \text{DET}}{\partial z} + \frac{\partial}{\partial z} \left(K_v(z) \frac{\partial \text{DET}}{\partial z} \right) \tag{29}$$

where W_d (m d⁻¹) is the detritus sinking rate.

2.3.4 Phytoplankton biomass, ~~trait distribution,~~ and primary production

405 The Eulerian module ~~additionally calculates the~~ also calculates the concentration of total phytoplankton nitrogen, carbon and chlorophyll ~~concentration, and biomass, as well as~~ the net primary productivity at each grid layer.

The concentration of total phytoplankton nitrogen, carbon and chlorophyll biomass (B_* ; B_{P_N} , mmol N m⁻³; B_{P_C} mmol C m⁻³; and $B_{P_{CHL}}$ mg Chl m⁻³) is computed at each grid layer as follows (note that we drop the subscript of z for convenience):

$$410 \quad B_* = \frac{\sum_{i=1}^k P_{i_{phy}}^* \cdot n_{i_{phy}}}{H}, \quad (30)$$

where ~~$P_{i_{phy}}^*$ refers to~~ $P_{i_{phy}}^*$ represents the three different units of mass (i.e., P_N , P_C and P_{CHL}).

Net Primary Productivity (~~NPP~~ NPP , mg C m⁻³ d⁻¹) of a vertical layer is integrated over a daily cycle:

$$\text{NPP} = \frac{1}{H} \sum_{t=0}^{24} \sum_{i=1}^{k(t)} \left[\Delta P_{C,i}(t) n_i(t) \right] \quad (31)$$

where $k(t)$ is the number of phytoplankton super-individuals at time t .

415 2.4 Size Spectra

~~Size spectra have been widely used to provide useful insights into the size structure and energy flow of aquatic communities (Platt and Denman, 1977). To illustrate how this model can be used to simulate the plankton size distribution, we plotted the size spectra for both phyto- and zooplankton at surface water and compared them between winter (31st March) and summer (31st August).~~

420 ~~To compute the size spectra, we counted the abundances of phytoplankton and zooplankton into the 20 size bins designed for zooplankton, although the phytoplankton cells were missing in large size classes.~~

~~Zooplankton abundances were estimated as the ratio between the zooplankton carbon divided by the individual carbon content of each size class. The individual carbon content of zooplankton was estimated from their biovolume following the allometric relationship proposed by Harris et al. (2000). First, based on the ESD of each zooplankton size class, their cell~~
425 ~~volume was calculated:~~

$$Z_C = \rho_Z \frac{\pi}{6} \text{ESD}^3 p_d p_c,$$

~~where ρ_Z is the density of zooplankton organisms, assumed to be equal to seawater density (1.025 g cm⁻³). p_d (dimensionless) is the proportion of wet mass constituted by dry mass, and p_c (dimensionless) is the proportion of dry mass constituted by carbon. We used the p_d and p_c values defined for non-gelatinous zooplankton (0.20 and 0.45, respectively). The carbon content~~
430 ~~of each zooplankton size class was converted from zooplankton nitrogen using the Redfield ratio (C:N = 106:16).~~

~~The size spectra of both phytoplankton and zooplankton were calculated as Ordinary Least-Squares (OLS) regression lines with log10(abundance) as the response variable and log10(biovolume) as the predictor. All zero abundance data were removed.~~

In addition, the abundance data of the two smallest size classes ($0.8\text{ }\mu\text{m}$, $1.2\text{ }\mu\text{m}$) were removed for zooplankton because such small size classes, deviating from the normal linear trend, were not often considered in the construction of size spectrain that layer.

2.4 Physical forcing

As a case study, our 1D model simulates the upper 250 m of the Bermuda Atlantic Time-series Study (BATS) station. To maintain generality and simplicity, some region-specific details, although some details (e.g., such as phosphorus limitation instead of nitrogen limitation) are not considered to keep its generality. A total, are not included. The vertical grid consists of 100 depth levels define the vertical grid, with increasing resolution towards levels, with resolution increasing toward the sea surface, following a sigma grid, approach similar to that used in the Regional Ocean Modeling System (ROMS; Shchepetkin and McWilliams (2005)).

The model is forced with three external externally forced by three environmental variables: temperature, vertical eddy diffusivity (K_v), and photosynthetically active radiation (PAR). Temperature profiles are imported from (T , $^{\circ}\text{C}$) are prescribed from monthly climatological data in the World Ocean Atlas 2013 (Locarnini et al., 2024). Vertical eddy diffusivity profiles (K_v , $\text{m}^2\text{ s}^{-1}$) were interpolated from is interpolated from daily outputs of a previous model output (Bruggeman and Bolding, 2014; Le Gland et al., 2021). We define the Mixed Depth Layer (MLD, m) as the depth at which simulation using a 1D General Ocean Turbulence Model (GOTM) which uses the $k \sim \epsilon$ turbulence model to compute $K_v < 10^{-4}\text{ m}^2\text{ s}^{-1}$.

Surface (Bolding and Villarreal, 1999; Bruggeman and Bolding, 2014; Le Gland et al., 2021). Finally, surface PAR (I_0 , W m^{-2}) at BATS station was estimated is estimated dynamically at each time step as a function of mid-day light, time of the day, and day length, following Anderson et al. (2015). Light values along attenuation within the water column (I_z , W m^{-2}) are estimated is computed at each time step following using the Beer-Lambert law, based on I_0 and chlorophyll concentration Law, incorporating the effects of seawater and chlorophyll absorption:

$$I_z = I_0 e^{-z (K_w + K_{\text{Chl}} \int_z^0 \text{Chl}(z) dz)}, \quad (32)$$

where K_w (m^{-1}) and K_{Chl} ($\text{m}^2 (\text{mg Chl m}^{-2})^{-1}$) are the light attenuation coefficients for seawater and chlorophyll-a, respectively.

Fig. 2 shows the temporal and spatial distribution seasonal evolution of the physical forcing variables, representing the typical seasonal illustrating the typical temporal and spatial variability at the BATS station. During winter and early spring, the temperatures were low temperatures remain low, while K_v were is high. The mixed layer reaches a maximum depth of around Mixed Layer Depth (MLD, m), defined as the depth where $K_v < 10^{-4}\text{ m}^2\text{ s}^{-1}$, reaches its maximum depth ($\sim 250\text{ m}$) from February to the end of March. Starting from April, an abrupt decrease in MLD is observed, associated with a decrease late March. From April onward, MLD decreases abruptly, coinciding with a reduction in K_v and an increase in temperature in the first 25 m of the water column. From late spring to late a warming of the surface water. Throughout late spring and summer, the water column is strongly stratified. It can be observed how becomes strongly stratified, with the surface mixed layer gets

465 warmer and reaches a depth ~ warming and deepening to 50 m in by September. At the beginning-onset of fall, a decrease in temperature and an increase in cooling temperatures and increasing K_v deepens deepen the surface mixed layer once again.

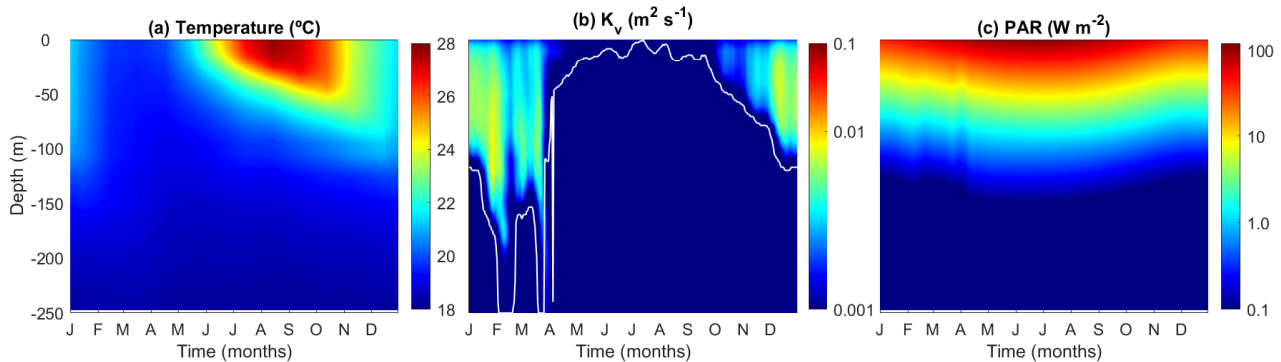


Figure 2. Temporal and vertical variability of the model forcing variables. (a) Temperature ($^{\circ}C$), (b) Vertical diffusivity (K_v , $m^2 s^{-1}$), and (c) Photosynthetically active radiation (PAR; $W m^{-2}$). In (b), the white line identifies the mixed layer depth (m).

2.5 Vertical movements transport of particles

The movement of both-

470 In addition to the super-individuals and passive particles due to diffusion is simulated as, the Lagrangian module also simulates the movement of passive inert particles as a validation tool. Both passive particles and super-individuals move due to vertical diffusion, following a random walk following model based on Visser (1997). The following equation computes the change in position (z_t) for-of an individual particle from the, from depth at time t to time $t + 1$ (z_{t+1}), is computed over a finite time step δt as:

$$z_{t+1} = z_t + K'_v(z_t) \delta t + R \sqrt{2 r^{-1} K_v(z_t + \frac{1}{2} K'_v(z_t) \delta t) \delta t}, \quad (33)$$

475 where $K'_v(z_t)$ represents the gradient of diffusivity vertical diffusion gradient ($= \delta K_v / \delta z$) at depth z_t , R is a random factor corresponding to a uniformly distributed random variable with a mean of zero variable drawn from a uniform distribution with zero mean and variance r ($r = \frac{1}{3}$ for a uniform distribution between - 1 and 1) (Ross and Sharples, 2004).

We-In addition to diffusion, phytoplankton vertical movement is also influenced by sinking, which is size-dependent and follows the allometric relationship described in Durante et al. (2019) (Table 2). However, we assume no vertical velocity-of
480 currents-current velocity in the system. The sinking rate-of-

2.6 Initial conditions

In our standard model run, we initialize 20,000 phytoplankton cells depends on cell size, following the allometric relationship in Durante et al. (2019) super-individuals and 1,000 passive particles with these numbers maintaining constant during the

simulation. The vertical positions of both passive particles and phytoplankton super-individuals were randomly assigned
485 between the surface (0 m) and bottom (250 m) at the start of the simulation, following a uniform distribution.

As previously explained, the three master traits were assigned to each phytoplankton super-individual at the beginning of
the simulation. The initial ESD of each phytoplankton super-individual was randomly drawn from a uniform distribution in log
space between 0.80 and 60.00 μm . These ESD values were then converted to phytoplankton cellular carbon content following
Menden-Deuer and Lessard (2000) (Table 2). Subsequently, the initial cellular nitrogen content was estimated based on the
490 Redfield ratio (C:N = 106:16 mol:mol), while the initial cellular chlorophyll content was calculated assuming a Chl:C ratio
of 1:50 (g:g). C_{div} , the maximum carbon content at which a division event occurs, was set twice the initial cellular carbon
content. The initial number of phytoplankton cells per super-individual was calculated assuming a constant phytoplankton
nitrogen concentration of 0.10 mmol m^{-3} throughout the water column.

~~To ensure that~~ T_{opt} values were randomly drawn from a uniform distribution between 2 and 30 °C. Similarly, α^{Chl} values
495 were randomly drawn from a uniform distribution in log space between 0.01 and 0.50 ($\text{W}^2 \text{m}^{-1} \text{g Chl (mol C)}^{-1} \text{d}^{-1}$).

In the Eulerian module, the initial concentrations of DIN were interpolated from the January profile at BATS, obtained from
the World Ocean Atlas (Garcia et al., 2024). The total initial zooplankton nitrogen concentration was set to 0.10 mmol N m^{-3}
throughout the water column and evenly distributed among the 20 size classes. Detritus concentration was also initialized as
0.10 mmol N m^{-3} .

500 2.7 Boundary conditions

To preserve total nitrogen, zero fluxes are applied at both surface and bottom boundaries for all Eulerian fields. Additionally,
a reflective boundary condition is used for particles encountering both surface and bottom boundaries during the random
walk. ~~works correctly, the Lagrangian module also computes the movements of 1000 passive inert particles whose trajectories
can be tracked.~~ We leave the option to use the Dirichlet boundary condition if one wishes to.

505 2.8 Model simulations

To achieve (quasi) steady-state seasonal cycles, PIBM was run for 6 years, with only the last year analyzed in this study.
The model solves the differential equations governing biological processes using the forward Euler method with a constant
10-minute time step. However, because the vertical random walk requires a shorter time step (Ross and Sharples, 2004),
particles undergo 200 substeps per biological time step, resulting in a random walk time step of 3 seconds.

510 As the particle random walk is computationally intensive, we implemented parallel computing using openMPI (Gabriel et al., 2004)

Table 1. Fixed parameters of the ~~1D~~-~~hybrid~~-1D PIBM model.

Parameter	Symbol	Value	Units
Phytoplankton maximal chlorophyll to nitrogen ratio ¹	θ_{\max}^N	3	g Chl mol N ⁻¹
Phytoplankton cost of biosynthesis ¹	ζ	3.00	mol C mol N ⁻¹
Shape-factor describing the dependence of V_{\max}^C on Q^{N1}	n	1.00	dimensionless
Phytoplankton maximal growth rate when $T_{\text{opt}} = 15\text{ }^{\circ}\text{C}^2$	μ'	5	d ⁻¹
Interspecific activation energy ²	E_i	0.22	eV
Normalization constant for E_d^2	E_{d0}	2.3	eV
Normalization constant for E_a^2	E_{a0}	0.98	eV
Activation energy for heterotrophic processes ³	E_z	0.65	eV
Boltzmann constant	k_b	8.62×10^{-5}	eV K ⁻¹
Scaling exponent against θ for E_a^2	β	-0.20	eV
Scaling exponent against θ for E_d^2	ϕ	0.27	eV
Normalization constant of the σ_{PSII} and θ^C relationship ⁴	δ	0.492	m ² W ⁻¹ g C g Chl ⁻¹
Exponent of the σ_{PSII} and θ^C relationship ⁴	κ	0.469	dimensionless
Turnover time of the electron transfer chain ⁴	τ	5.50×10^{-3}	s
Damage constants of a PSU ⁴	k_d	5.00×10^{-6}	dimensionless
<u>Mutation probability of a phytoplankton cell</u>	<u>ν_0</u>	<u>10^{-12}</u>	<u>cell⁻¹</u>
<u>Standard deviation of mutation of three traits</u>	<u>σ</u>	<u>0.1</u>	<u>dimensionless</u>
Zooplankton grazing half-saturation constant	$K_{P,j_{\text{pred}}}$	0.15	mmol N m ⁻³
Coefficient of the prey refuge ⁵	Λ	-6.60	(mmol N) ⁻¹ m ³
Zooplankton gross growth efficiency ⁶	ξ	0.30	dimensionless
Fraction of unassimilated food by zooplankton ⁶	η	0.24	dimensionless
Zooplankton mortality rate at 15 °C	m_z	0.005	d ⁻¹
Standard deviation of zooplankton grazing kernel ⁵	$\sigma_{j_{\text{pred}}}$	0.5	dimensionless
Conversion rate from detritus to DIN at 15 °C	R_{dn}	0.10	d ⁻¹
Detritus sinking rate	W_d	0.50	m d ⁻¹
Light attenuation coefficient for seawater ⁷	K_w	0.04	m ⁻¹
Light attenuation coefficient for chlorophyll ⁷	K_{Chl}	0.025	(mg Chl m ⁻²) ⁻¹

1. Geider et al. (1998); 2. Chen (2022); 3. Brown et al. (2004); 4. Nikolaou et al. (2016); 5. Ward et al. (2012); 6. Buitenhuis et al. (2010); 7. Gan et al. (2010).

Table 2. Size scaling coefficients ~~of~~ for phytoplankton and zooplankton traits, following the general formula ($y = aV^b$; $y = aV^b$), where V represents cell volume (μm^3). For phytoplankton, V is derived from C_{div} using the allometric relationship from Marañón et al. (2013) (first entry in this table). For the size scaling of maximal phytoplankton growth rate and respiration rates, see Eqs. 13 and 14.

Variable	Symbol	a	b	Units
Phytoplankton cellular carbon ¹	P_C	0.20 <u>0.017</u>	0.88	pmol C cell ⁻¹
Phytoplankton half-saturation constant for nitrogen uptake ²	K_N	0.14	0.33	mmol N m ⁻³
Phytoplankton cellular maximum N:C ratio ³	Q_{max}^N	0.25	-0.07	mol:mol
Phytoplankton cellular minimal N:C ratio ^{4,3}	Q_{min}^N	0.07	-0.17	mol:mol
Sinking rate of phytoplankton ^{5,4}	W_{phy}	0.0019	0.43	d ⁻¹
Zooplankton maximum grazing rate ^{6,3}	$G_{max} I_{max}$	21.90	-0.16	d ⁻¹

1. ~~Menden-Deuer and Lessard (2000)~~ Marañón et al. (2013); 2. Edwards et al. (2012); 3. ~~Marañón et al. (2013)~~; 4. ~~Ward et al. (2012)~~; 5. ~~Durante et al. (2019)~~ Ward et al. (2012); 6. ~~Durante et al. (2019)~~.

2.9 Initial conditions Observational data

Initial concentrations of DIN were interpolated from the January profile of BATS from the World Ocean Atlas (Garcia et al., 2024). The total initial concentration of all zooplankton size classes was assumed as 0.1 mol N m^{-3} throughout the water column and was split equally among 20 size classes. We initialized 20000 phytoplankton super-individuals and 1000 passive particles and these numbers do not change during the simulation. The vertical positions of both the passive and the phytoplankton For model validation, we obtained observational data for DIN (nitrate and nitrite), Chl, and NPP from the BATS website (<https://bats.bios.asu.edu/>). To interpolate the data for each vertical grid at each time point, we applied the k -nearest neighbors (KNN) algorithm, with each data point calculated as the mean of the three nearest neighbors.

520 2.10 Mean trait and trait (co)variance

We characterize phytoplankton community composition in terms of mean trait values and trait (co)variance. Trait (co)variance serves as a measure of functional diversity, capturing trait variations in the community (Norberg et al., 2001; Chen et al., 2019; Le Gland et al., 2020). The community mean phytoplankton trait is calculated as the carbon-weighted mean of all phytoplankton cells in the community:

$$525 \quad \bar{l} = \frac{\sum_{i=1}^k l_i n_i P_{C,i}}{\sum_{i=1}^k n_i P_{C,i}} \quad (34)$$

where l_i and n_i represent the trait value and number of cells of super-individual ~~particles were randomly assigned between the surface~~ (i , respectively, and $P_{C,i}$ represents the cellular carbon content of this super-individual.

The trait covariance is calculated as:

$$\text{COV}(l_j, l_m) = \frac{\sum_{i=1}^k (l_{j,i} - \bar{l}_j) (l_{m,i} - \bar{l}_m) n_i P_{C,i}}{\sum_{i=1}^k n_i P_{C,i}} \quad (35)$$

530 where \bar{l}_j and \bar{l}_m represent the mean values of trait j and m , respectively. $l_{j,i}$ and $l_{m,i}$ represent the trait j and m of super-individual i respectively. It is important to note that we treat $\log(C_{div})$, T_{opt} , and $\log(\alpha^{chl})$ as traits as they are more likely to follow normal distribution than the raw units.

2.11 Functional diversity: Rao index

535 We computed the functional diversity of a phytoplankton community using the Rao index, a classic index to quantify diversity in functional ecology (Rao, 1982; De Bello et al., 2021):

$$Rao = \sum_{i=1}^N \sum_{j=1}^N p_i p_j d_{ij} \quad (36)$$

where p_i and p_j represent the relative proportion in carbon biomass of super-individual i and j , respectively, in the community. d_{ij} represents the Euclidean distance between these two super-individuals in the three-dimensional trait space:

$$d_{ij} = \sqrt{\sum_{k=1}^3 (l'_{i,k} - l'_{j,k})^2} \quad (37)$$

540 where $l'_{i,k}$ represents the k^{th} trait of super-individual i normalized between 0 m) and the bottom (250 m) at the start of the simulation following a uniform distribution and 1 because the three traits have different units ($l'_{i,k} = \frac{l_{i,k} - l_{min}}{l_{max} - l_{min}}$).

The ESD of each phytoplankton As shown above, the Rao index not only captures the relative abundance (biomass in this case) of each individual, but also takes into account trait similarities between individuals.

2.12 Size Spectra

545 Size spectra have been widely used to provide insights into the size structure and energy flow of aquatic communities (Platt and Denman, 1990). To illustrate how this model can be used to simulate the plankton size distribution, we plotted the size spectra for both phyto- and zooplankton at the first 10 m of the water column and compared them between winter (20th February) and summer (20st August). We constructed the size-abundance spectra by calculating the total cell abundances in each of the 20 size bins that cover the entire size range of the simulated plankton community (from ~ 0.80 to $3600 \mu\text{m}$). The size bins were defined according to an octave (\log_2) scale of cell volume.

The size-abundance spectra for phytoplankton is computed straightforwardly, with cellular volume derived using the allometric relationship between volume and cellular carbon (C_{div} , see Table 2). Each super-individual was randomly assigned between

0.80 and 60.00 μm following a uniform distribution on the log space. The phytoplankton cellular carbon content was derived from cell volume following Mara $\tilde{\text{n}}$ on et al. (2013) (Table 2). The initial cellular nitrogen content was then estimated following the is assigned to the corresponding size bin based on its cellular volume, and the total abundance within each size bin is determined by summing the number of phytoplankton cells within all super-individuals of that volume range.

For zooplankton, the biovolume of each of the 20 size classes is known, but their abundance must be estimated from biomass concentration. Specifically, the abundance of each size class is estimated as the ratio of total zooplankton carbon biomass to the individual carbon content. To calculate individual carbon content, zooplankton nitrogen content is first converted to carbon for each size class using the canonical Redfield ratio (C:N = 106:16 mol:mol). The initial cellular chlorophyll content was estimated assuming a Chl:C ratio of 1:50 (g:g). The initial number of cells per phytoplankton super-individual was calculated assuming a uniform concentration of phytoplankton nitrogen of 0.1 mmol m⁻³. Individual carbon content (Z_C , g C) is then derived from biovolume using the allometric relationship proposed by Harris et al. (2000):

$$Z_C = \rho_Z \frac{\pi}{6} \text{ESD}^3 p_d p_c, \quad (38)$$

where ρ_Z (g cm⁻³) is the density of zooplankton, assumed to be equal to seawater density (1.025 g cm⁻³) throughout the water column.

2.13 Boundary conditions

To preserve total nitrogen, we apply Neumann boundary conditions for both surface and bottom boundaries. We also assume a reflective boundary for particles that encounter both surface and bottom boundaries during random walk. We leave the option to use the Dirichlet boundary condition if one wishes to. p_d (dimensionless) is the proportion of wet mass composed of dry mass; and p_c (dimensionless) is the proportion of dry mass composed of carbon. For non-gelatinous zooplankton, we adopted values of $p_d = 0.20$ and $p_c = 0.45$.

2.13 Model simulations

To reach (quasi)-steady-state seasonal cycles, the size spectra of both phytoplankton and zooplankton were calculated as ordinary least-squares regression lines with log₁₀-transformed total cell abundance (y-axis, individuals m⁻³) as the response variable, and log₁₀-transformed cell volume for each size class (x-axis, μm^3) as the predictor. All zero-abundance data were removed. Additionally, the abundance data of the 1D-hybrid model was run for 6 years. Only the last year is analyzed in this study. We use the forward Euler method with a constant time step of 10 min throughout to numerically solve the differential equations of biological processes. However, because the vertical random walk requires a short time step (Ross and Sharples, 2004), we allowed the particles to have 100 time steps within each biological time step (i.e., a time step of 6 s for the random walk). Two smallest size classes (0.80 μm , 1.20 μm) were removed for the zooplankton fraction because such small size classes, which deviate from the normal linear trend, were not often considered in the construction of size spectra.

As the particle random walk is a significant step consuming computing time, we implemented parallel computing to run particle random walk using open MPI (Gabriel et al., 2004).

585 2.13 Sensitivity analysis

2.14 ~~Observational data~~

~~For validating-~~

2.13.1 Number of phytoplankton super-individuals

590 The number of phytoplankton super-individuals is a crucial parameter affecting the computation speed of PIBM. Ideally, we aim to determine the minimum number of super-individuals that can be used without compromising model accuracy. To assess the impact of the number of phytoplankton super-individuals on model outputs, we ~~downloaded observational data of DIN (nitrate and nitrite), Chl, and NPP from the BATS website (-). To interpolate the data of each vertical grid at each date from the irregular cruise data, we used *k-nearest-neighbours* (KNN) algorithm with each data point calculated as the mean of the three nearest neighbours~~ ran simulations with 5,000, 10,000, and 50,000 super-individuals while keeping all other parameters constant and compared their results against the standard run with 20,000 super-individuals.

2.13.2 Comparison with simple NPZD models

To examine the effects of increased model complexity proposed in the PIBM, and assess how incorporating multiple phytoplankton traits influences plankton biomass and production, we implemented two simple NPZD models: one Eulerian and one Lagrangian. These models contain only a single, generic phytoplankton group, where carbon, nitrogen, and chlorophyll dynamics follow Geider et al. (1998) (Eqs. 1, 2, 3).

However, important differences exist. First, the term of photoinhibition is not included, meaning that P^C is an increasing function of light (I) in both models:

$$\sim \alpha^{Chl} \theta^C IP_m^C$$

(39)

605

Second, μ_m is an exponential function of temperature following the Arrhenius equation:

$$\mu_m = \mu_0 e^{E_p x}$$

(40)

where E_p is the activation energy of phytoplankton.

Table 3. Parameters of the simple NPZD model that differ from PIBM.

Parameter	Symbol	Value	Units
Phytoplankton maximal growth rate at 15 °C	μ_0	2	d ⁻¹
Activation energy of phytoplankton growth	E_p	0.32 ¹	eV
Initial slope of the P-I curve	α^{Chl}	1.2	m ² W ⁻¹ g C (g Chl) ⁻¹
Zooplankton maximal ingestion rate at 15 °C	I^{max}	1.6	d ⁻¹
Zooplankton grazing half-saturation constant	K_P	0.15	mmol N m ⁻³
Zooplankton mortality rate at 15 °C	m_z	0.1	d ⁻¹

1. Chen et al. (2012).

Last, there is only one zooplankton group that feeds exclusively on phytoplankton and experiences linear mortality:

610
$$\frac{\partial ZOO}{\partial t} = f_h(T) \xi ZOO I^{max} \frac{P_N}{P_N + K_P} (1 - e^{\Lambda P_N}) - ZOO m_z f_h(T) + \frac{\partial}{\partial z} \left(K_v(z) \frac{\partial ZOO_j}{\partial z} \right) \quad (41)$$

We ran these two simple NPZD models under the same forcing as PIBM with identical simulation settings (e.g., time step, simulation duration). However, we have to stress that strict comparisons between PIBM and simpler NPZD models are impossible due to different parameter sets and model architecture.

3 Results

615 Below, we first describe the behavior of the phytoplankton model, followed by ~~an in-depth examination of the 1D model output.~~
~~Then we compare the modelled~~ a detailed analysis of PIBM outputs. We then compare the modeled fields of DIN, Chl, and NPP
against observations to ensure that our model can ~~at least~~ qualitatively reproduce the main ~~patterns of observations~~ observed
patterns. Afterwards, we ~~show some patterns~~ highlight patterns emerging from the model ~~outputs~~ that may be interesting but
hard to directly measure *in situ*. Finally, we present the results of sensitivity analyses, including the effects of phytoplankton
620 super-individuals and comparisons against simple NPZD models.

3.1 Phytoplankton fitness landscapes

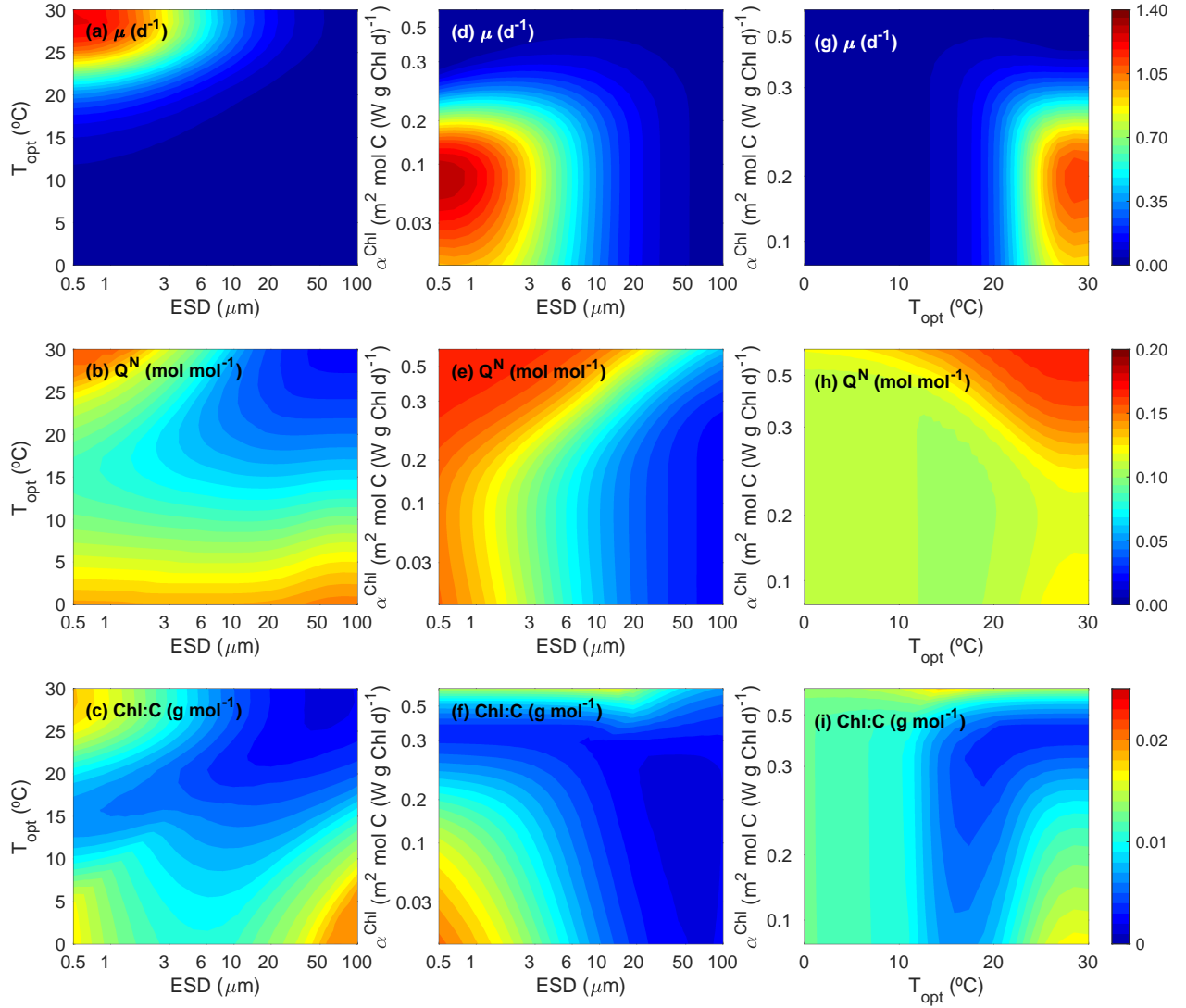


Figure 3. Contour plots for phytoplankton growth rate (μ , d^{-1}), N:C ratio (Q^{N} , mol:mol) and Chl:C ratio (g:g) as functions of the three master traits (size, ESD (μm); optimal temperature, T_{opt} ($^{\circ}\text{C}$); and light affinity, α^{Chl} ($\text{W}^2 \text{m}^{-1} \text{g Chl mol C}^{-1} \text{d}^{-1}$)) at equilibrium under a typical summer condition (dissolved inorganic nitrogen = $0.10 \text{ mmol N m}^{-3}$, temperature = 28°C , and PAR = 250 W m^{-2}).

Fig. 3 illustrates how growth rate (μ , d^{-1}) (top row), Q^{N} :N:C (mol:mol) (middle row), and the Chl:C ratio (g:g) (bottom row) vary for different phytoplankton ecotypes which have unique combinations of traits (i.e., ~~a phytoplankton with a characteristic~~

ESD-calculated-from C_{div} , T_{opt} and α^{Chl}). To facilitate understanding the responses, we plot the interpretation, we present 2D contour plots by varying two traits that vary two traits at a time (i.e., ESD vs. T_{opt} , ESD vs. α^{Chl} , and T_{opt} vs. α^{Chl}), while keeping the third trait constant. Below, we present Figure 3 shows the results for the summer period, characterised by a constant DIN concentration of 0.10 mmol m⁻³, a temperature of 28°C, and an irradiance of 250 W m⁻².

In the left column, we present the effects of impact of cell size and T_{opt} , with a constant α^{Chl} of 0.08 W² m⁻¹ g Chl⁻¹ mol C d⁻¹. This α^{Chl} value is designed to adapt to the current light conditions without strong photoinhibition. As anticipated, the maximum peak of μ is observed for small-size phytoplankton (< 3 μ m), efficient at taking up nitrogen at low environmental concentrations, and with a T_{opt} around 28-30 °C, close to the environmental temperature (Fig. 3a). Moreover, we observe close to zero μ values for the whole size range when T_{opt} is far from the environmental temperature, highlighting maladapted ecotypes. Q^N :N:C shows maximum values for small phytoplankton cells, almost regardless of T_{opt} (Fig. 3b), decreasing towards largest sizes, consistent with the observations by Baer et al. (2017). In this particular scenario, where both irradiance and α^{Chl} are constant, the Chl:C ratio also peaks for small cells with high T_{opt} (Fig. 3c), similar to μ . The increase in the Chl:C ratio for large-size ecotypes and T_{opt} < 10 °C may be related to the minimum growth in carbon content (Fig. 3a).

In the middle column, we present how growth, Q^N :N:C and Chl:C ratios change with cell size and α^{Chl} , with a constant T_{opt} of 30 °C. This T_{opt} value is close to the environmental temperature (28 °C), indicating adapting to the ambient temperature. μ peaks for small phytoplankton (< 3 μ m) and α^{Chl} around 0.10 W² m⁻¹ g Chl⁻¹ mol C d⁻¹ (Fig. 3d). The light condition of 250 W m⁻² favor organisms with low light affinity values (α^{Chl} < 0.20 W⁻¹ m² g Chl⁻¹ mol C d⁻¹) with high capability of photo-repair. Q^N :N:C mainly depends on cell size, progressively decreasing towards the largest sizes (Fig. 3e). Other things All else being equal, the cells with more optimal α^{Chl} values tend to have exhibit lower N:C ratios (Fig. 3e). Chl:C ratio shows the maximum values for the smallest phytoplankton with the lowest α^{Chl} , and the minimum values for the largest phytoplankton with the highest α^{Chl} (Fig. 3f). As larger α^{Chl} increases the probability of photoinhibition and reduces the carbon-specific photosynthesis rate, the amount of cellular chlorophyll content also decreases relative to carbon.

In the right column, we present the effects-impact of T_{opt} and α^{Chl} , with a constant ESD of 1.36 μ m. As anticipated, given the defined environmental conditions, the highest growth rate (μ) is achieved by phytoplankton cells with a T_{opt} close to the environmental temperature and α^{Chl} values around 0.10 W⁻¹ m² g Chl⁻¹ mol C d⁻¹, which are better adapted to high irradiance (Fig. 3g). Moreover, Q^N :N:C increases with T_{opt} , with a minimum corresponding to cells with high growth (Fig. 3h). Similarly, the Chl:C ratio is maximized for ecotypes with T_{opt} closer to the environmental temperature, and characterized by lower α^{Chl} values, resulting in reduced photoinhibition (Fig. 3i).

3.2 Comparisons with observations

We compared the observed DIN, Chl, and NPP at the BATS station with the model output (Fig. 4). Our model is able to reproduce the general qualitative patterns of these three variables, albeit with some quantitative differences. Both the observation and the model output show an increase in surface DIN and Chl during the winter when mixing is the strongest (Fig. 2). The model also successfully reproduces the deep chlorophyll maximum between 50 and 100 m. In spite of Despite the presence of the deep chlorophyll maximum, the model can also reproduce the surface maximum of NPP, which extends below 50 m during

the summer. In other words, our model manages to resolve the paradox of low Chl but high NPP at surface waters during the summer in the oligotrophic subtropical ocean.

Admittedly, there are some differences between the model output and the observations. The model appears to underestimate DIN in deep waters and overestimate Chl in the euphotic zone, which is likely because the phosphorus limitation at BATS is not considered in our model (Ammerman et al., 2003).

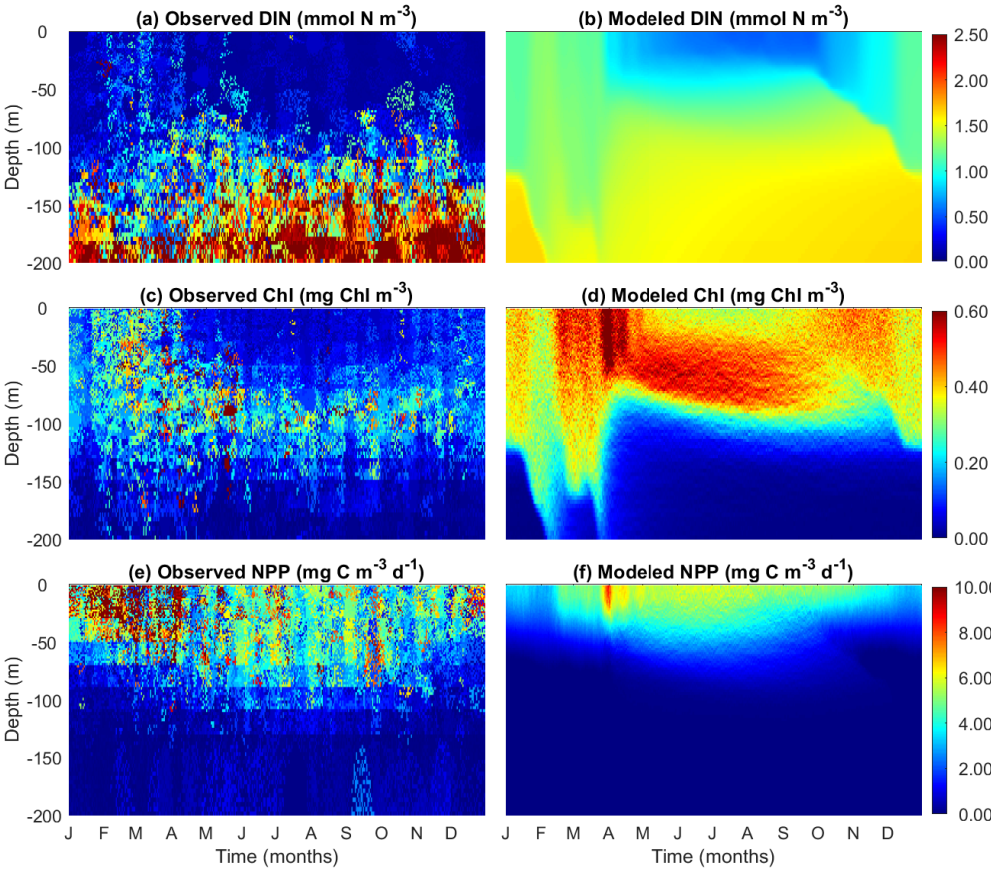


Figure 4. Comparing Daily comparison between observed data (left column) and modeled data (right column) for (a, b) ~~DIN~~ dissolved inorganic nitrogen (DIN, mmol m^{-3}), (c, d) ~~Chl~~ chlorophyll concentration (Chl, mg m^{-3}), and (e, f) ~~NPP~~ net primary productivity (NPP, $\text{mg C m}^{-3} \text{ d}^{-1}$) between observations and model outputs.

3.3 Modelled-Modeled patterns of passive particles

Fig. Figure 5 compares the vertical distribution of passive particles and phytoplankton super-individuals, enabling-allowing us to verify if whether the Lagrangian module properly-models-accurately represents the random walk of particles. While

the distribution of phytoplankton super-individuals can be additionally affected by cell division and death, the distribution of passive particles is driven purely by diffusion.

As passive particles (Fig. 5a) are homogeneously distributed throughout the water column, we can affirm that the particle random walks are working correctly. During periods of high vertical mixing, the distribution of super-individuals shows a homogeneous pattern (Fig. 5b). However, as expected, during the period of strong stratification of the water column (April - October), more pronounced differences were observed. A greater number of super-individuals were found in the euphotic layer, where conditions for survival are optimal, with their abundance decreasing in deeper layers.

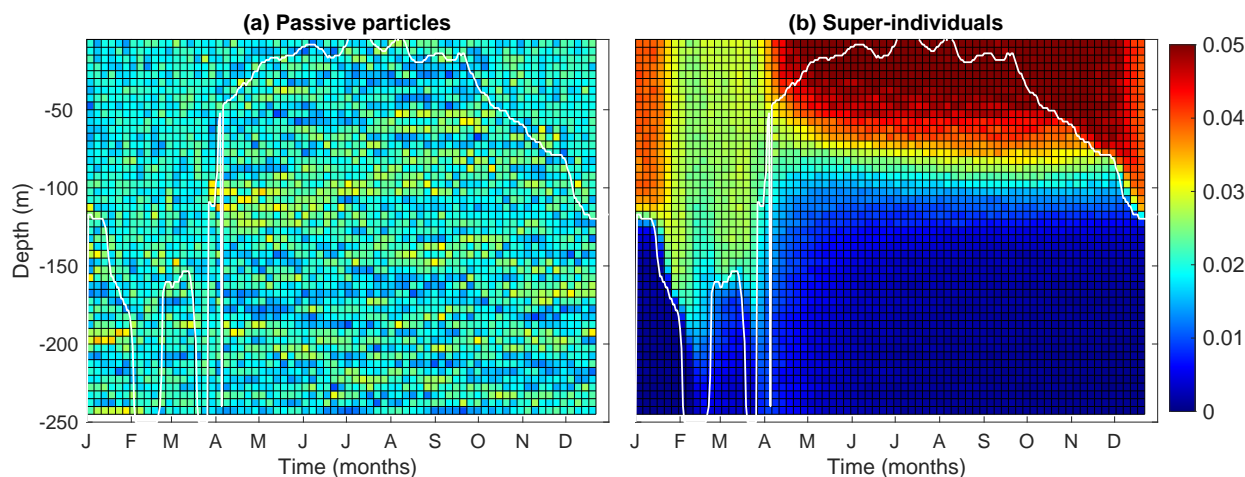


Figure 5. Mean temporal and vertical frequencies of (a) passive particles and (b) phytoplankton super-individuals every 5 m depth and every 5 days along the last year's simulation. The white lines represent the ~~Mixed Layer Depth~~ mixed layer depth (MLD, m).

3.4 Modelled patterns of phytoplankton carbon, nitrogen, zooplankton and detritus

Phytoplankton carbon and nitrogen concentrations, for which no observational data are available, show similar patterns. Both of them peak at spring after the shoaling of mixed layer depth (Fig. 6A,Ba,b). Before the spring bloom, surface phytoplankton carbon and nitrogen can penetrate deeper than 150 m due to strong winter mixing. Similar to the patterns of chlorophyll, phytoplankton carbon also shows elevated concentrations at the deep chlorophyll maximum during summer, although this maximal layer is not as pronounced as that of chlorophyll (Fig. 4d).

By contrast, zooplankton biomass peaks in the summer and autumn after the phytoplankton spring bloom (Fig. 6Cc), suggesting that the demise of the phytoplankton spring bloom is at least attributed to intensified zooplankton grazing. Detritus also peaks after the phytoplankton spring bloom, but unlike zooplankton, its concentration gradually declines with time, as a result of both accelerated decomposition due to high temperature and shifts in zooplankton size structure.

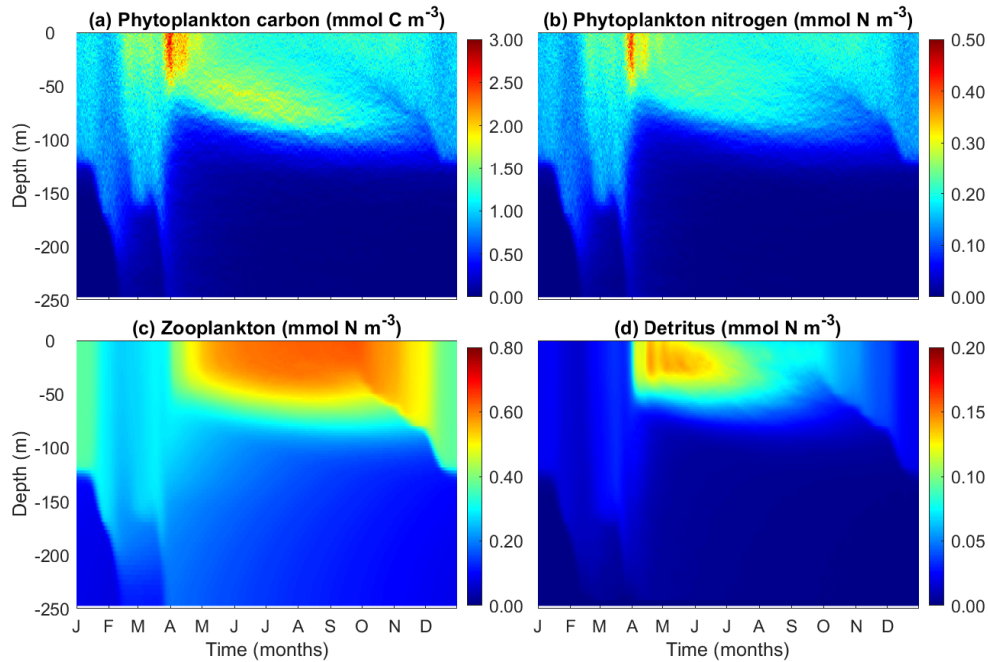


Figure 6. Modelled patterns—Temporal and vertical variability of modeled (a) phytoplankton carbon (mmol C m^{-3}), (b) phytoplankton nitrogen (mmol N m^{-3}), (c) total zooplankton nitrogen (mmol N m^{-3}), and (d) detritus (mmol N m^{-3}).

3.5 Modelled Chl:C and N:C ratios

The phytoplankton cellular N:C and the Chl:C ratios are important phytoplankton properties that show the extent key indicators of how phytoplankton cells acclimate to the acclimate to light and nutrient environment. They availability. These ratios are also crucial to link for linking phytoplankton carbon to satellite observations of chlorophyll to satellite-derived chlorophyll observations and to the limiting element nutrient, nitrogen (Fig. 7).

The model output provides—predicts higher Chl:C ratios in the surface mixed layer during the winter as a result of winter due to enhanced nutrient supply due to from strong mixing. The In contrast, Chl:C ratios are low—lower in both the surface layer in the summer summer surface layer and in deep waters beneath the surface below the mixed layer. These patterns can be understood as an outcome of the balance between photosynthesis and chlorophyll synthesis. Under high light and low nutrient conditions, phytoplankton cells reduce the rate of chlorophyll synthesis relative to carbon synthesis and vice versa. However, when the light is too low, the synthesis rate of chlorophyll is too low to sustain the maintenance of chlorophyll, leading to a phenomenon known as “bleaching” (Pahlow et al., 2013; Behrenfeld et al., 2016).

While the patterns of Chl:C ratios can be easily understood from the perspective of environmental control, N:C ratios are more related to changes in phytoplankton size than to the environment DIN or light. While one might expect that phytoplankton cells should have lower N:C at the surface during the summer due to low DIN, the model actually predicts the opposite pattern.

This is because the summer surface waters are dominated by small cells which tend to have larger N:C ratios (Marañón et al., 2013; Baer et al., 2017).

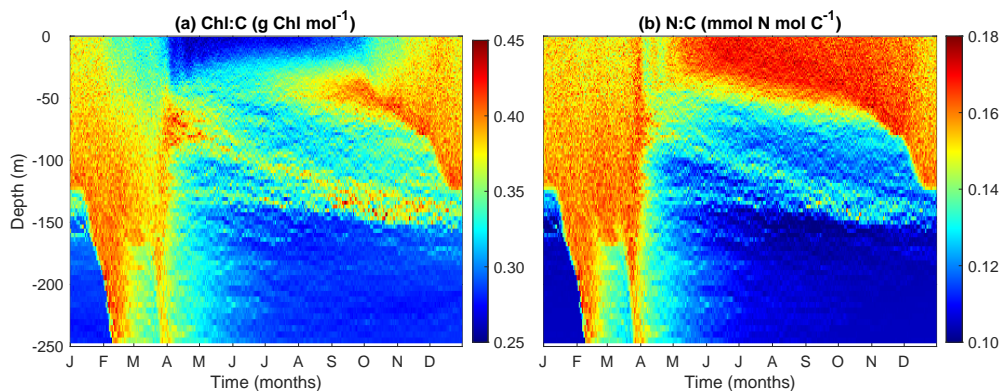


Figure 7. ~~Modelled patterns~~ Temporal and vertical variability of ~~modeled~~ phytoplankton (a) Chl:C ratio (g Chl mol C^{-1}) and (b) N:C ratio (mol N mol C^{-1}).

700 3.6 Modelled phytoplankton trait distribution and functional diversity

Our model ~~shows a typical pattern that~~ reproduces the typical pattern of increasing phytoplankton mean size ~~increases~~ with nutrient availability (Fig. 8Aa). Phytoplankton mean size is the smallest at the surface during the summer and autumn, where DIN is low, but increases with depth as nutrients become more abundant. During winter, phytoplankton mean size is also larger at the surface during the winter (but not as large as in deeper waters during summer) when nutrient concentrations are higher
705 due to stronger mixing.

Phytoplankton size variance ($\text{Var}(C_{div})$) is an index for size diversity and is the greatest in the area of the deep chlorophyll maximum during summer and autumn and the lowest beneath the euphotic zone (150 m) (Fig. 8Bb). The high size variance at the deep chlorophyll maximum results from the movements of small cells from above and large cells from below. Size variance is also slightly greater at the surface during the winter than the summer due to more abundant nutrients, but not as high as
710 those in the deep chlorophyll maximum. This suggests that dispersal probably plays the most important role in affecting size diversity.

As expected, phytoplankton community mean T_{opt} largely follows seawater temperature, with higher values at the surface during summer and lower values at depth (Fig. 8Cc). However, phytoplankton mean T_{opt} ~~show~~ shows the lowest values at the area of deep chlorophyll maximum during summer and autumn, corresponding to the maximal variance of T_{opt} (Fig. 8Dd).
715 Otherwise, the variance of T_{opt} generally shows larger values in the surface mixed layer than at depth.

Phytoplankton mean light affinity (α^{Chl}) also follows the opposite pattern of light availability, being the lowest at the summer surface and the highest in the deepest waters (Fig. 8Ee). The variance of α^{Chl} shows a qualitatively similar pattern with those

of variance of C_{div} and T_{opt} , being higher during the winter and at the deep chlorophyll maximum (Fig. 8Ff), again suggesting mixing can enhance trait variance and diversity.

720 The covariances between traits suggest selection forces against different traits. The covariance between T_{opt} and C_{div} is largely negative during the winter, suggesting larger cells tend to be cold-adapted at local scales. Conversely, $Cov(T_{opt}, C_{div})$ is often positive at surface and deep chlorophyll maximum during summer at local scales. It is important to note that these covariances are calculated at local scales. We can still observe negative $Cov(T_{opt}, C_{div})$ for the whole water column in which the small and warm-adapted cells are at the surface and large and cold-adapted cells are at depth.

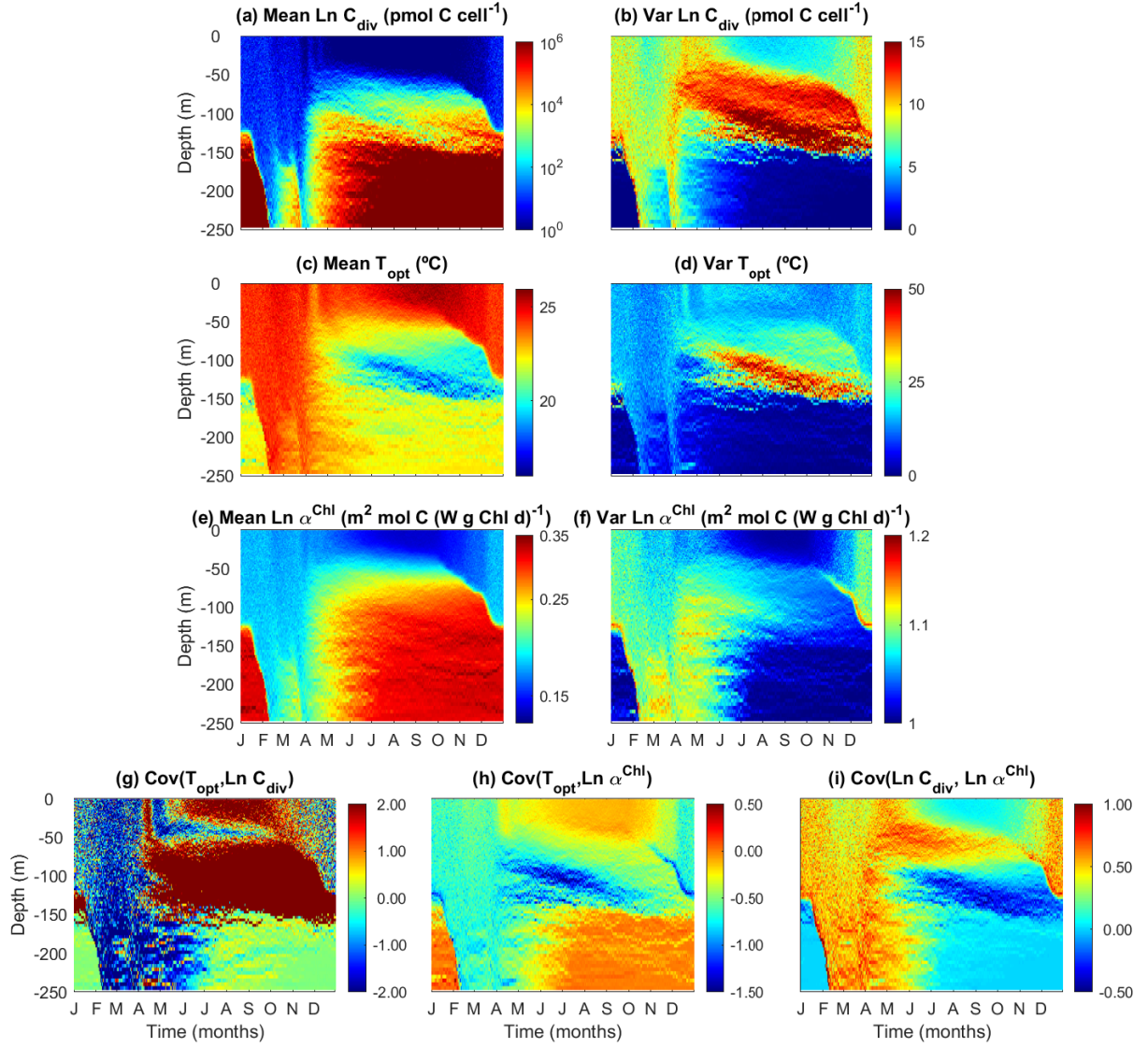


Figure 8. Temporal and vertical variability of modeled phytoplankton traits weighted by phytoplankton carbon content (Eq. 34). (a) Mean phytoplankton size (ESD) back-transformed from the carbon threshold of cell division (C_{div}). (b) Variance of phytoplankton C_{div} . (c) Mean phytoplankton optimal temperature (T_{opt}). (d) Variance of phytoplankton T_{opt} . (e) Mean phytoplankton light affinity represented by ln-transformed slope of Photosynthesis-Irradiance curve (α^{Chl}). (f) Variance of phytoplankton α^{Chl} . (g) Covariance between phytoplankton T_{opt} and $\ln C_{div}$. (h) Covariance between phytoplankton T_{opt} and $\ln \alpha^{Chl}$. (i) Covariance between phytoplankton $\ln C_{div}$ and $\ln \alpha^{Chl}$.

725 The covariance between T_{opt} and α^{Chl} is largely negative, suggesting that due to the negative environmental correlation between temperature and light, cold-adapted cells tend to adapt to the low light. The covariance between size and α^{Chl} is mostly positive during winter and at deep chlorophyll maximum, indicating larger cells tend to adapt to low light.

~~Temporal and vertical patterns of traits weighted by phytoplankton carbon content (Eq. 34). (A) Mean phytoplankton size (ESd) back-transformed from the carbon threshold of cell division (C_{div}). (B) Variance of phytoplankton C_{div} . (C) Mean phytoplankton optimal temperature (T_{opt}). (D) Variance of phytoplankton T_{opt} . (E) Mean phytoplankton light affinity represented by ln-transformed slope of Photosynthesis-Irradiance curve (α^{Chl}). (F) Variance of phytoplankton $\ln \alpha^{Chl}$. (G) Covariance between phytoplankton T_{opt} and $\ln C_{div}$. (H) Covariance between phytoplankton T_{opt} and $\ln \alpha^{Chl}$. (G) Covariance between phytoplankton $\ln C_{div}$ and $\ln \alpha^{Chl}$.~~

730 The modelled patterns of functional ~~richness largely follow those of (eo)variances~~ diversity quantified as the Rao index largely follow the trait variances and the covariance between cell size and other two traits (Fig. ??). Both functional richness and evenness are 9). The Rao index is greater during the winter season with strong mixing and at the layer of deep chlorophyll maximum.

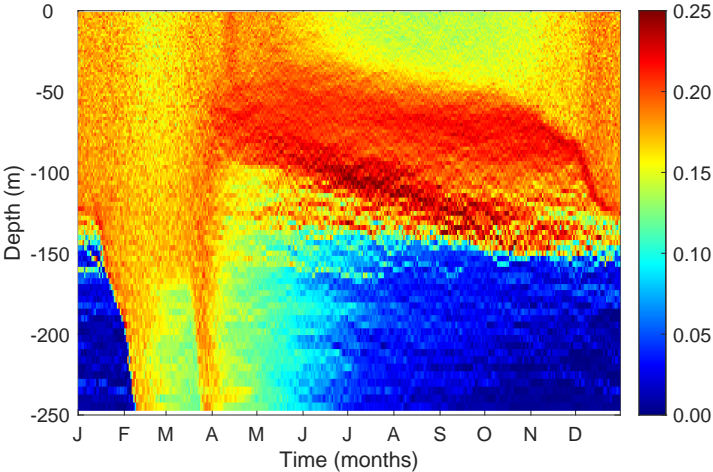


Figure 9. ~~Modelled patterns~~ Temporal and vertical variability of phytoplankton functional diversity (a) richness and (b) Rao index evenness weighted by cellular carbon.

3.7 Plankton size spectra

During both seasons, log abundances form linear relationships with log size (biovolume) for both phytoplankton and zoo-
 740 plankton (Fig. 10). The slopes of phytoplankton size spectra were between ~~-1.35 and -1.1~~ -0.99 and -1.36, consistent with what would be expected for phytoplankton communities in oligotrophic oceans (Fig. 10) (Marañón, 2019). The slopes of the phytoplankton size spectra were more negative in the summer than in the winter, also consistent with previous observations that the

phytoplankton size spectra became steeper and small phytoplankton became more dominant when nutrient supply diminishes (Huete-Ortega et al., 2014).

745 By contrast, the slopes of zooplankton were similar between summer ~~and winter~~ (-0.85) and winter (-0.80) and were less steep than those of phytoplankton (Fig. 10). This relates to three factors. First, large zooplankton have a wider feeding kernel than small ones, thus having access to a wide range of prey. Second, the predator-prey size ratio increases with predator size, which makes the slope of size spectra less steep (Trebilco et al., 2013). Third, large zooplankton can also feed on small zooplankton. These patterns are consistent with the observations in marine plankton communities that more biomass can exist

750 in larger size classes, which leads to an inverted biomass pyramid (Gasol et al., 1997; Trebilco et al., 2013).

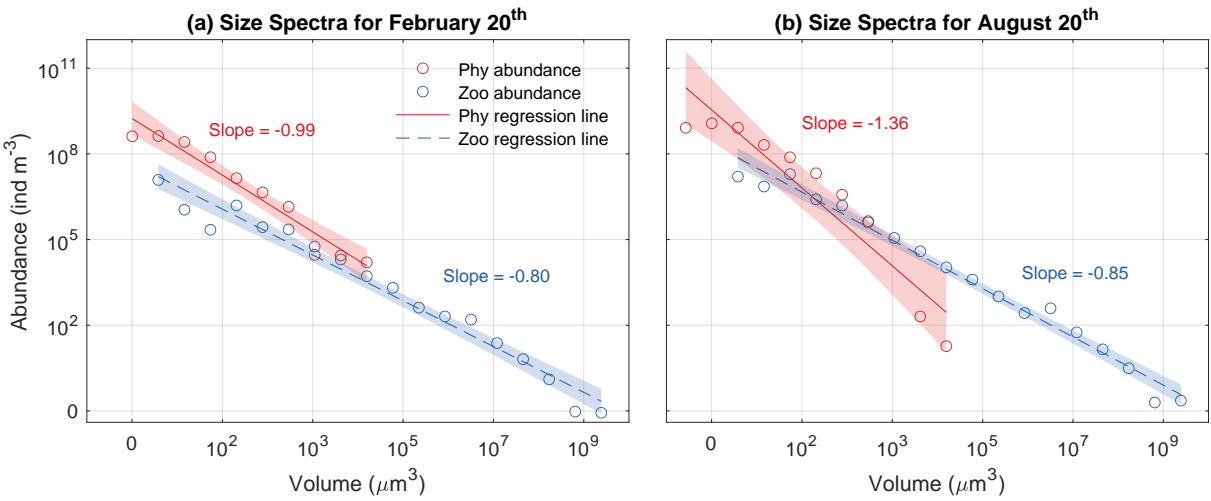


Figure 10. Phytoplankton (Phy) and zooplankton (Zoo) size spectra at the surface during ~~the~~ (a) ~~Winter~~ winter (31st March 20th February) and (b) ~~Summer period~~ summer (31st–20th August). ~~The~~ Open dots ~~show~~ represent the abundance of each size class. ~~The~~ straight, ~~while~~ solid lines ~~show~~ indicate the OLS ordinary least squares regression ~~lines~~ fits. Shaded areas denote the 95% confidence intervals ($\alpha = 0.05$). The slopes of each regression line are also ~~shown~~ displayed.

3.8 Diel variability of phytoplankton cell

Our model allows to track the properties of a phytoplankton cell throughout a diel cycle to gain insights from the life history of the cell. ~~Fig.~~ Figure 11 shows the trajectory of a randomly selected super-individual at hourly resolution during the first week of ~~the~~ winter and the first week of ~~the summer period (hourly resolution)~~ summer.

755 During the winter period, this phytoplankton particle was dispersed widely between the surface and 50 m in the water column due to strong mixing (Fig. 11a). Because the water column was well mixed, it was exposed to relatively stable conditions of temperature (~ 19.50 °C) and DIN (Fig. 11b,d). By contrast, this particle experienced variable PAR conditions, oscillating between daily maximums ~ 50 ~~W m⁻²~~ and 200 W m^{-2} . By comparison, the variations of Chl:C ratio (θ^C) are modest (Fig. 11h), suggesting its limited acclimation capability.

760 During the summer period, the super-individual position was more stable at around 35 m in the first seven days due to weak mixing (Fig. 11a). As a consequence, the environmental conditions (temperature, DIN, and PAR) the particle experienced were relatively stable. However, the variabilities of phytoplankton Q^N and θ^C N:C and Chl:C are of the same magnitude between summer and winter, mainly due to diel changes of light.

765 The phytoplankton Q^N and θ^C The phytoplankton N:C and Chl:C were higher during the winter period (Fig. 11g, red line) than during the summer (blue line), reflecting the higher nutrient and lower light environment in the winter.

In spite of Despite the seasonal differences, both particles divided twice during this period (Fig. 11e-f). The phytoplankton cellular carbon content increased progressively with time before the division event until reaching the division threshold (Fig. 11f). When this threshold was reached, the number of phytoplankton cells doubled and the cellular carbon, nitrogen and Chl content halved (Fig. 11e-f).

770 Irrespective of the birth event, we can also observe clear diel changes in cellular carbon, nitrogen, and chlorophyll contents induced by light-driven photosynthesis, nutrient uptake, chlorophyll synthesis, and dark respiration. Cellular carbon increased from sunrise to sunset but declined in the dark due to respiration. Correspondingly, N:C and Chl:C ratios declined during the daytime and increased during nighttime.

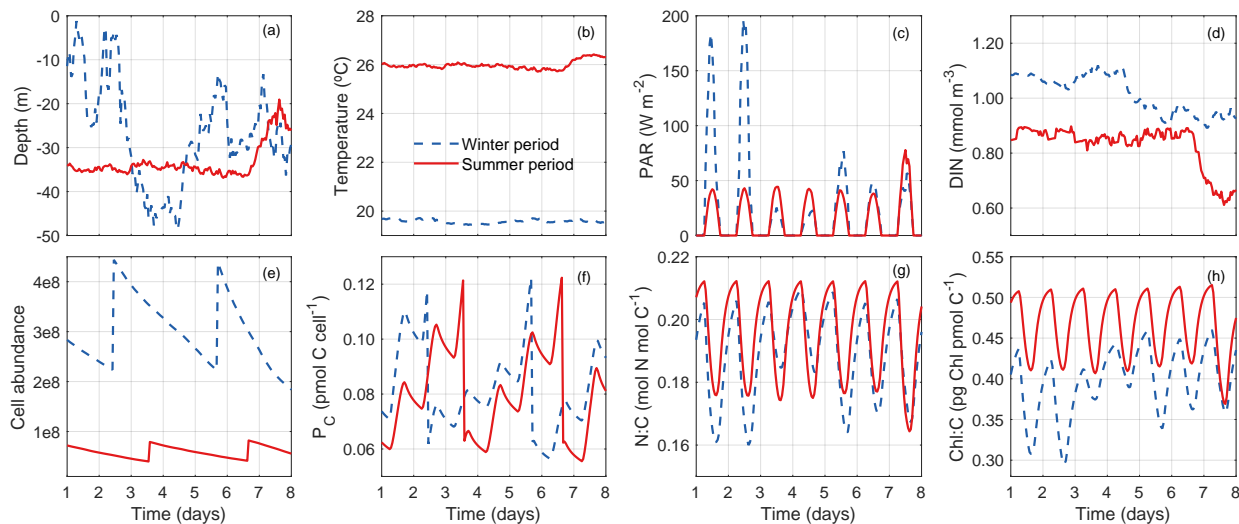


Figure 11. Tracking a randomly selected phytoplankton super-individual for 7 days (hourly resolution) during the winter (red-solid-blue dashed lines) and the summer (blue-dashed-red solid lines) periods. (a) Depths (m) between which the super-individual oscillated in the water column. (b) Differences in the Temperature (°C) conditions experienced by the super-individual. (c) Differences in the Photosynthetically Active Radiation conditions (PAR; $W m^{-2}$) experienced by the super-individual. (d) Differences in the nitrogen concentration conditions (NO_3 , $pmol N m^{-3}$) experienced by the super-individual. (e) Temporal evolution of phytoplankton cell abundance (number of cells) within the selected super-individual. (f) Temporal evolution of cellular carbon content (P_C , $pmol C cell^{-1}$). (g) Temporal evolution of nitrogen cellular quota (Q^N , $pmol N pmol C^{-1}$). (h) Temporal evolution of chlorophyll to carbon cellular ratio (θ^C , $mg Chl mmol C^{-1}$). Vertical dashed lines indicate the beginning of each day. (The panel labels are missing):

3.9 Effect of the number of super-individuals

We investigated the effect of the number of phytoplankton super-individuals on the key model outputs such as including DIN, phytoplankton biomass, NPP, and diversity by running three-four simulations with 5,000, 10,000, and 20,000, and 50,000 super-individuals. We computed the temporal trends of key-these variables integrated throughout the water column during the final year of each simulation (Fig. ??12).

The number of phytoplankton super-individuals has negligible effects on the bulk properties such as DIN, ZOO, P_N , P_C , Chl and NPP (Fig. ??a-f). However, the number of phytoplankton super-individuals has noticeable effects on phytoplankton functional richness and evenness. The lower numbers of weak effects on these variables. Increasing the number of super-individuals lead to lower estimates of phytoplankton functional richness and evenness than tends to reduce DIN and increase phytoplankton biomass and NPP, although the trends are very similar between 20,000 and 50,000 super-individuals during most times of the year. Moreover, the simulation with lower numbers of. Contrary to what one might expect, increasing the number of super-individuals struggles to converge to regular seasonal cycles of functional richness and evenness, suggesting that a sufficient number of phytoplankton slightly reduces phytoplankton functional diversity particularly in the winter and summer, although the seasonal trends are qualitatively similar. In summary, the number of phytoplankton super-individuals is required to represent phytoplankton diversity does not have an appreciable effect on modelled phytoplankton biomass and diversity.

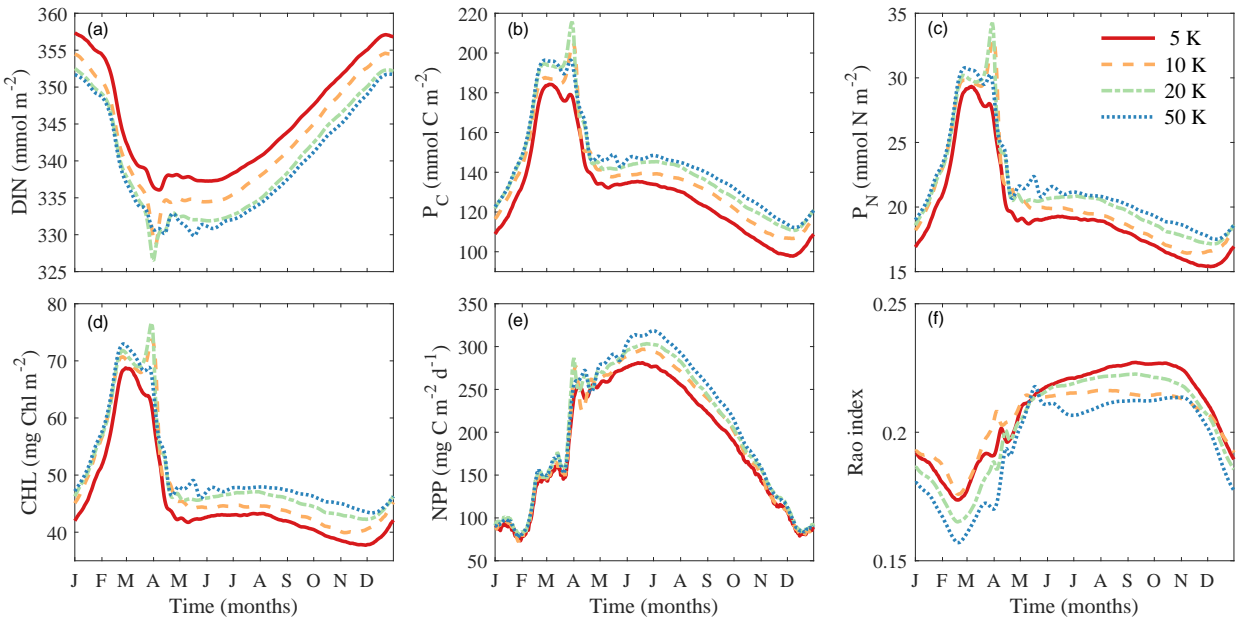


Figure 12. Comparisons of simulations with different numbers of super-individuals on vertically integrated variables. (a) Dissolved inorganic nitrogen (DIN, mmol m^{-2}). (b) Phytoplankton carbon (P_C , mmol C m^{-2}). (c) Phytoplankton nitrogen (P_N , mmol N m^{-2}). (d) Chlorophyll (Chl, mg Chl m^{-2}). (e) Net primary production (NPP, $\text{mg C m}^{-2} \text{d}^{-1}$). (f) Phytoplankton functional diversity (Rao index).

3.10 Comparisons with simple NPZD models

790 We compared the model outputs of DIN, Chl, and NPP between two simple NPZD models, Eulerian and Lagrangian, and
PIBM (Fig. 13) and the three models produce qualitatively similar seasonal and vertical patterns. The three models generate
similar concentrations of DIN, although DIN is depleted earlier during spring in PIBM than in the two NPZD models. The two
NPZD models also reproduce deep chlorophyll maximum layers during summer and autumn, similar to PIBM. A pronounced
difference is that the Chl concentrations are lower and the deep chlorophyll maximum is deeper in the two NPZD models than
795 PIBM. The two NPZD models also produce higher NPP levels than PIBM, although the seasonal and vertical patterns are
qualitatively similar.

Notably, the outputs produced by the Lagrangian NPZD model are almost identical with those produced by the Eulerian
NPZD model, although with slightly greater spatiotemporal variations.

In summary, the outputs generated by PIBM are not qualitatively different from simpler NPZD models, although some
800 details differ due to model architecture and parameter setting.

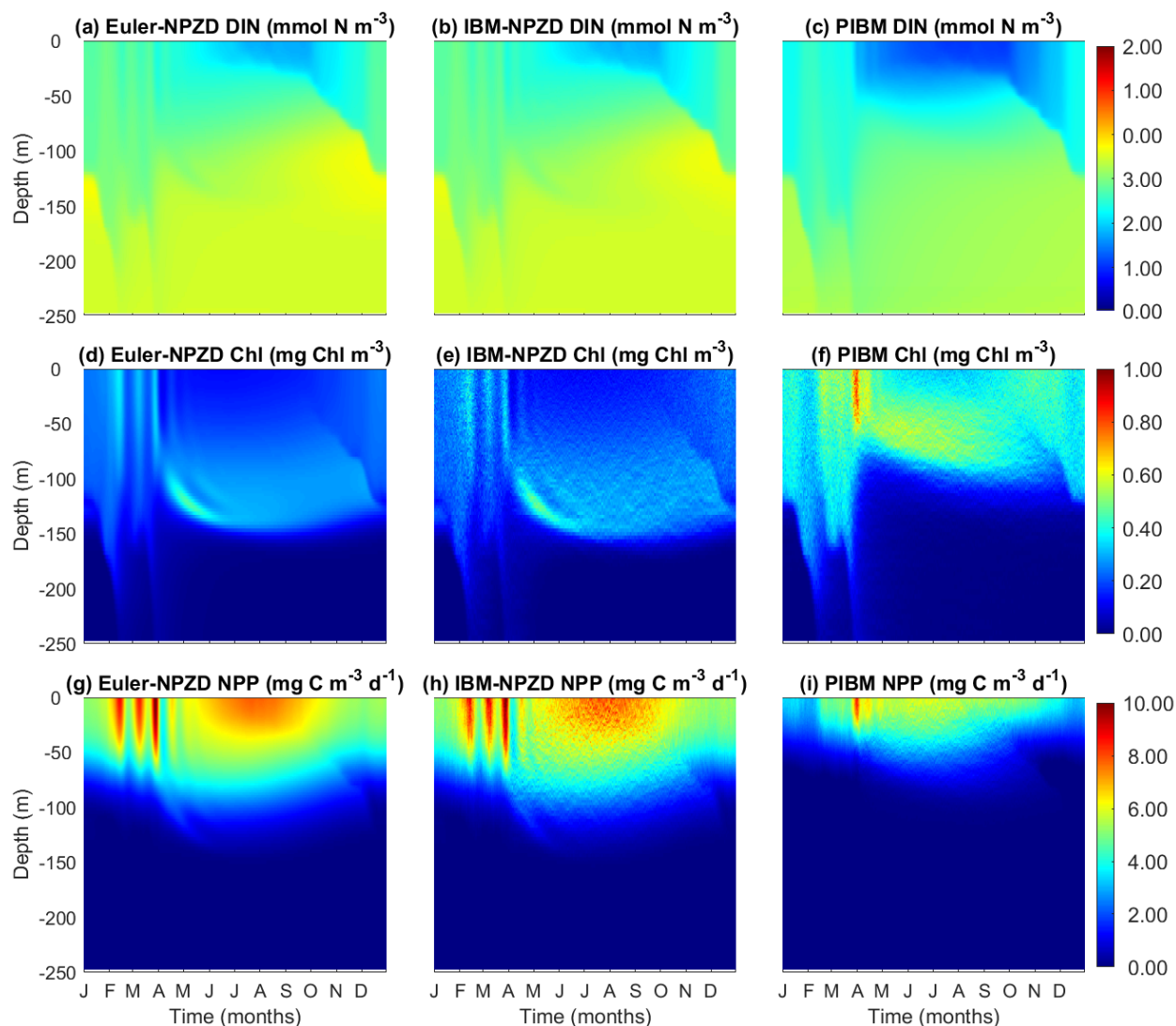


Figure 13. Comparisons Temporal and vertical variability comparison of simulations with different numbers the output results of super-individuals on vertically-integrated variables. the three models: the Eulerian NPZD model (a left column) Dissolved inorganic nitrogen (DIN, mmol N m^{-3}), the Lagrangian NPZD model (b middle column) Phytoplankton carbon (P_C , mmol C m^{-3}) and the PIBM (right column). (a, b, c) Phytoplankton nitrogen Dissolved Inorganic Nitrogen (P_N DIN, mmol N m^{-3}). (d, e, f) Chlorophyll a (Chl, mg m^{-3}). (g, h, i) Net primary production (NPP, $\text{mg C m}^{-3} \text{d}^{-1}$). (f) Zooplankton biomass (ZOO, mmol N m^{-3}). (g) Phytoplankton functional richness. (h) Phytoplankton functional evenness.

4 Discussion

We have presented a novel hybrid Eulerian-Lagrangian plankton model (PIBM) that treats phytoplankton as particles or ~~superindividuals. We assign super-individuals. Each phytoplankton super-individual is associated with~~ three master traits (~~cell size- cell size~~, (C_{div}) , temperature affinity (T_{opt}), and light affinity (α^{Chl}) ~~to each phytoplankton super-individual and these~~
805 ~~master traits determine many other physiological traits involved in- which govern various physiological traits related to~~ nutrient uptake and photosynthesis. ~~We also incorporate acclimation and evolution in this model~~ PIBM also enables phytoplankton cells to have acclimate and evolve, making it ~~suitable- well-suited~~ for addressing many questions in phytoplankton ecology and evolution.

4.1 A brief history of individual-based models on phytoplankton

810 ~~In recent decades, individual-based models have gained popularity in ecology modelling. This approach allows to simulate the non-linear effects of the physical environment (Woods and Onken, 1982), the interaction between individuals, the life cycles of individuals (Grimm et al., 2006; Hellweger et al., 2008; Hense and Beckmann, 2010), the adaptive behaviour, and the intrapopulation variability in response to internal and external environmental conditions (Hellweger et al., 2008; Grimm et al., 2010)~~
-

815 ~~In oceanography, the first individual-based models were developed between the late 1970s and early 1980s (Ledbetter, 1979; Platt and Ga-~~
- ~~However, because of their high computational cost, it was not until recent decades that they started to gain momentum (Cianelli et al., 2004; Woods, 2005; Nogueira et al., 2006; Cianelli et al., 2009; Beckmann et al., 2019; Ranjbar et al., 2021). To obtain a realistic representation of the phytoplankton community~~ Before discussing the merits and limitations of PIBM, it is necessary to model a large number of Lagrangian particles. Each particle is modelled individually, thus obtaining a unique
820 ~~history of interaction with the environment and other particles. Therefore, the computational cost will be proportional to the number of particles considered. One way to limit this computational cost is to model Lagrangian particles as super-individuals (Scheffer et al., 1995), although this approach may restrict the population's heterogeneity and diversity (Ranjbar et al., 2021) if the total number of particles is not high enough. With the increasing computing power, this problem is of less severity in individual-based models.~~

825 ~~Several studies, employing the Lagrangian Ensemble model, have also depicted zooplankton as Lagrangian particles (agents) (Woods and Barkmann, 1993, 1994; Woods, 2005; Nogueira et al., 2006). However, these studies define a singular class of herbivorous zooplankton that feeds indiscriminately on all phytoplankton particles, irrespective of their size. Furthermore, they do not account for the interaction between a phytoplankton particle and zooplankton, essentially adopting an Eulerian approach. In other words, when calculating grazing, a zooplankton particle is allowed to consume phytoplankton particles~~
830 ~~within the same grid layer without considering physical interaction. While our model has considered zooplankton size structure and size-dependent feeding kernels, it will be our next step to model zooplankton as Lagrangian particles which can account for more detailed interactions between zooplankton and phytoplankton particles.~~

In summary, individual-based models are gaining popularity thanks to the increasing computing power and they can provide novel insights that conventional models fail to provide.

835 4.1 Strengths and limitations of the model

There are several strengths of our model. First, the individual-based phytoplankton model captures the realistic physical movement and acclimation status of a phytoplankton cell affected by ocean turbulence. It has been widely acknowledged that for both Eulerian models and *in situ* incubations, the estimates of primary production can be biased although the extent of this bias is uncertain (Barkmann and Woods, 1996; Ross and Geider, 2009; Baudry et al., 2018). Part of the differences in the bias estimates may be due to the traits of the phytoplankton cells. Most of the previous individual-based phytoplankton models do not consider phytoplankton functional diversity (i.e., the different traits associated with each cell). While the aim of this paper is not to compare the primary production estimates between the Eulerian model and Lagrangian model (as developing an Eulerian version of the same plankton model is beyond the scope of this paper), it remains to be investigated how the estimates of primary production can differ between the Eulerian model and Lagrangian model in different environments (e.g., stratified vs. well-mixed water column).

Second, our phytoplankton model captures three dimensions of phytoplankton traits. We have followed the DARWIN model (Follows et al., 2007; Barton et al., 2010; Dutkiewicz et al., 2020) to assign three master traits to phytoplankton albeit with slight differences. The assumption of designing these three traits is that they are largely orthogonal to each other. In other words, small and large cells have an equal probability of being warm-adapted or cold-adapted and they also have the same probability of being high-light-adapted or low-light-adapted. How the changes of these traits affect phytoplankton acclimation and the differences in primary production estimates between Eulerian and Lagrangian models remain to be investigated using our model.

Third, our phytoplankton model allows phytoplankton evolution. We have built the functionality of phytoplankton evolution in the Lagrangian model enabling the mutation of all three master traits of phytoplankton super-individuals. Modelling phytoplankton evolution has been a hot topic recently (Ward et al., 2019; Beckmann et al., 2019) and the individual-based model is an ideal and straightforward approach to accommodate mutation and evolution (Acevedo-Trejos et al., 2022). We will use this model to further explore how evolution affects phytoplankton diversity and primary productivity. important to reiterate that its development was not aimed at improving the modeling of commonly measured bulk variables such as nutrient and chlorophyll concentrations. On the contrary, it is well known that more complex models often provide less accurate predictions of these variables than simpler models (Kwiatkowski et al., 2014), due to greater uncertainties in model structure and parameterization. Indeed, our simple NPZD models can generate outputs as good as, or even better than, PIBM without parameter tuning. However, this does not imply that developing complex models is unnecessary or without value. The same applies to this work, our model is designed as a tool to address research questions related to phytoplankton biodiversity, which simple NPZD models cannot address.

865 ~~We also~~

4.1 Model limitations

In this section, we highlight several limitations of ~~our model~~ PIBM which will be addressed in future work. ~~One weakness of the current model is the slow computation. To obtain a realistic representation of the phytoplankton community, it is necessary to model~~

870 A key limitation of PIBM is its high computational cost. This arises from the need for a short time step to accurately simulate the random walk of particles in the water column (Ross and Sharples, 2004). To balance this requirement (<6 s) with computational efficiency, we set the biological reaction time step to 200 times that of the random walk. Additionally, accurately representing the phytoplankton community requires modelling a large number of ~~particles. As super-individuals. Since each particle is modelled individually, the computational cost will be proportional to simulated individually, computational cost~~
875 scales proportionally with the number of particles ~~considered. To obtain a satisfactory model result. Our sensitivity analyses indicate that 5,000 super-individuals appear sufficient to reproduce the main seasonal patterns of phytoplankton diversity, it is desirable to model as many particles as possible and other variables, although some quantitative differences remain (Fig. ??), thus incurring heavy computation costs.~~

~~The short time step is needed for correctly simulating the random walk (Ross and Sharples, 2004). In fact, to meet the~~
880 ~~requirement of short timestep (<6 s) in the random walk but to minimize computation time, we have made the timestep of biological reactions 100 times longer than the time step of the random walk (12).~~

We ~~have~~ also implemented openMPI parallel computing for simulating the random walk of both the passive and phytoplankton particles. However, the computation is still ~~too slow~~ not fast enough to allow effective sensitivity analysis or parameter optimization. In the future, we will make the computations of biological reactions in parallel and will attempt openMP which may
885 allow more efficient memory sharing among threads. Another ~~more advanced~~ remedy would be to implement GPU computing to speed up the computation. We are considering to integrate our model into PlanktonIndividuals.jl (Wu and Forget, 2022), a Julia language-based phytoplankton individual-based model that can take advantage of GPU computing capability of the Julia language.

The second limitation relates to the inadequate parameterization of the phytoplankton model using laboratory data and the
890 need for more extensive validation of the overall ~~4D~~ model output. While the size scaling of nutrient uptake has been studied extensively in the literature (~~Edwards et al., 2012; Marañón et al., 2013~~) (e.g. Edwards et al., 2012; Marañón et al., 2013) and the temperature dependence of phytoplankton growth seems clear (Thomas et al., 2012; Chen, 2022), ~~how the light traits of phytoplankton affect growth are relatively unknown particularly when considering the effect~~ the influence of phytoplankton light traits on growth remains relatively unknown, particularly in the context of photoacclimation. Edwards et al. (2015)
895 ~~analysed~~ analyzed the relationships between light traits (e.g., slope of the P-I photosynthesis-irradiance curve). However, ~~the model they used did not consider their model did not account for~~ the dependence of photosynthetic rate on the Chl:C ratio or photoinhibition. To ~~consider~~ account for both the effects of the Chl:C ratio and photoinhibition on ~~photosynthesis of phytoplankton~~ phytoplankton photosynthesis, we have to assume an empirical relationship between α^{Chl} and the repair rate based on ~~the contrast differences~~

900 et al., 1998). However, it remains unclear how this scaling relationship ~~can be applied~~ applies to phytoplankton in general. ~~We need more data on this to obtain~~ More data are needed to establish a more reliable relationship between ~~the~~ photosynthetic parameters.

The model outputs of phytoplankton traits also need to be validated against observations. While the measurements of phytoplankton size structure can be obtained, other traits such as T_{opt} and α^{Chl} are difficult to measure *in situ* on a cellular basis. 905 Moreover, even ~~the~~ bulk properties such as the N:C and Chl:C ratios of the whole phytoplankton assemblage are ~~difficult~~ challenging to measure *in situ* due to the difficulties ~~associated with measuring in quantifying~~ phytoplankton carbon and nitrogen. Only a few studies have managed to measure cellular carbon and/or nitrogen of small size phytoplankton using flow cytometric sorting (Graff et al., 2012; Baer et al., 2017), ~~while, for~~ For larger cells, most studies relied on microscopic counting to estimate phytoplankton cell volume which can then be converted to carbon (Cloern, 2018) without any measurement 910 of cellular nitrogen. ~~These types~~ This type of information is essential for studying biogeochemical cycling and validating ecosystem model outputs.

Another limitation is that the model may be overly complicated if we want to understand the key factor in controlling some ecological phenomenon (*see* below). ~~The model has included~~ PIBM considers multiple traits and processes ~~which form an intertwined~~, forming an interconnected feedback network that makes it ~~challenging~~ difficult to isolate the direct effect of a 915 single factor. For example, if we aim to assess whether primary production is limited by nutrient supply or light availability by performing a single-factor perturbation experiment, the increase in nutrient supply will not only affect the nutrient status but also the trait distribution of the whole phytoplankton community. This change in the mean ~~trait of phytoplankton~~ phytoplankton trait depends on the existing trait diversity ~~of the community~~ (within the community and the mutation rate of individual cells) ~~about which we have little information~~ (Acevedo-Trejos et al., 2015; Chen et al., 2019). If, both of which remain poorly 920 understood (Acevedo-Trejos et al., 2015; Chen et al., 2019). However, if a user already knows that the system can be simplified (e.g., there is little variability in phytoplankton thermal or light traits), ~~the user~~ they can modify the initial ~~condition~~ conditions and the mutation rate to remove the unnecessary trait variance. Thus, one can simplify this model to a single-trait (e.g., size) phytoplankton model if desired.

Finally, needless to say that our model only considers one limiting nutrient – nitrogen – without considering other important 925 elements such as phosphorus or silicate, ~~which probably leads to some discrepency~~. This omission likely contributes to discrepancies between observations and ~~our~~ model outputs at BATS, where phosphorus is ~~limiting~~ the limiting nutrient. It is up to the user to decide whether to add these nutrients to the model and also depends on the question being asked.

4.2 Model strengths

Our model has several strengths:

930 First, PIBM simulates the movements and acclimation status of phytoplankton cells with different traits in the ocean and helps address the uncertainty of modeling using Eulerian models and measuring NPP using fixed-depth incubations. It has been widely acknowledged that for both Eulerian models and *in situ* incubations, the estimates of primary production can be biased, although the extent of this bias is uncertain (Barkmann and Woods, 1996; Ross and Geider, 2009; Baudry et al., 2018). Some

935 differences in bias estimates may stem from the traits of phytoplankton cells. Although our preliminary comparison between the simple Eulerian and Lagrangian models shows no significant differences (Fig. 13), it remains to be tested whether incorporating additional traits will influence the divergence between the two approaches. Notably, most previous individual-based phytoplankton models are similar to our simple Lagrangian NPZD model which does not consider phytoplankton functional diversity (i.e., trait differences among cells). Although developing an Eulerian version of PIBM is beyond the scope of this paper, we can investigate how the estimates of primary production estimates differ between Eulerian and Lagrangian models across different environments (e.g., stratified vs. well-mixed water columns) using a simplified version of PIBM.

940 Second, PIBM captures three dimensions of phytoplankton traits. We followed the DARWIN model (Follows et al., 2007; Barton et al., 2017) to assign three master traits to phytoplankton, with slight differences. The rationale of designing these three traits is that they are largely orthogonal to each other. In other words, small and large cells have an equal probability of being warm-adapted or cold-adapted and they also have the same probability of being high-light-adapted or low-light-adapted. The impact of trait changes on phytoplankton acclimation and differences in primary production estimates between Eulerian and Lagrangian models remains to be investigated using PIBM.

950 Third, our phytoplankton model allows phytoplankton evolution. We incorporated phytoplankton evolution into the Lagrangian module, allowing mutation of all three master traits in phytoplankton super-individuals. Modeling phytoplankton evolution has been a hot topic recently (Ward et al., 2019; Beckmann et al., 2019), and individual-based models provide an ideal and straightforward approach to incorporate mutation and evolution (Acevedo-Trejos et al., 2022). We will use PIBM to further investigate the impact of evolution on phytoplankton diversity and primary productivity.

4.3 Potential applications

Below we discuss several potential applications of our model. Note that these are not exhaustive ~~but just serve as potential interesting directions~~ but highlight promising directions for future research.

955 4.3.1 Validating *in situ* measurements of primary production

~~An obvious~~ As discussed above, a key application of our model is to ~~check the assess~~ bias in *in situ* ~~measurements of primary production for which primary production measurements, where~~ incubation bottles are tethered at fixed depths ~~and the phytoplankton cells within the bottles experience different light environments from those being mixed throughout the, exposing phytoplankton to different light conditions than those experienced in a dynamically mixed~~ water column. While several studies have attempted to address this problem using Lagrangian phytoplankton models (Barkmann and Woods, 1996; Baudry et al., 2018; Tomkins et al., 2020), they often overlook phytoplankton traits. The most important phytoplankton trait for this problem ~~would be are~~ likely light-related traits (e.g., α^{chl}), which are rarely considered in ~~phytoplankton Lagrangian Lagrangian phytoplankton~~ models.

965 In addition, it is not only the trait itself but also the trait distribution that can matter for primary production. In other words, phytoplankton diversity and community composition ~~have to be taken into account to assess how accurate the in situ measurements of primary production are~~ is essential for accurately assessing *in situ* primary production measurements. Since

phytoplankton trait distribution is not static across time and space, the bias also depends on the phytoplankton community being sampled, so there is no guarantee that the bias can be easily extrapolated to other cases. The caveat is that we need to know the phytoplankton trait distributions (which are even harder to measure) to assess the accuracy of primary production estimates.

4.3.2 Understanding what controls phytoplankton diversity

The central theme of ecology revolves around understanding the factors that regulate biodiversity. Vellend and Agrawal (2010) presented a unified view of four processes controlling biodiversity: selection, dispersal, drift, and evolution. While many studies investigate what regulates phytoplankton diversity in the ocean (Barton et al., 2010; Vallina et al., 2014; Righetti et al., 2019; Dutkiewicz et al., 2020), few managed to examine the holistic effects of all four processes on biodiversity.

~~Our model~~ PIBM already incorporates the processes of selection, dispersal, and evolution and can be easily adapted to ~~integrate drift to address this gap~~ include drift. One challenge would be again about the computational costs if one wishes to understand the large-scale patterns of phytoplankton diversity by coupling the individual-based model with a global circulation model (Hellweger et al., 2014). Nevertheless, this challenge can be addressed using advanced computing techniques as described above.

4.3.3 Understanding how phytoplankton acclimation and trait distribution affect phytoplankton distribution

Biological oceanographers have long been fascinated by distinctive patterns of phytoplankton distribution in the ocean such as the deep chlorophyll maximum (Cullen, 2015) and spring bloom (Behrenfeld, 2010; Lévy, 2015). Despite extensive study, debates persist on what mechanisms drive these patterns.

The formation of deep chlorophyll maximum remains enigmatic with ongoing debate as to whether it reflects an actual accumulation of phytoplankton biomass or is primarily driven by photoacclimation for which phytoplankton cells increase their intracellular pigment content to acclimate to the low light condition. There is evidence that some phytoplankton species may be more abundant at the surface but show a peak of total pigments at the layer of deep chlorophyll maximum, while some other species may indeed show greater biomass at this layer (Chen et al., 2011). Again, the location of the phytoplankton biomass peak likely depends on phytoplankton traits. Our model is an ideal tool for elucidating the contribution of real biomass versus the photoacclimation effect to the layer of deep chlorophyll maximum.

Similar arguments can be raised for the spring bloom. Behrenfeld (2010) argued that contrary to the light effects (i.e., the critical depth hypothesis), the coupling between phytoplankton growth and zooplankton grazing plays a critical role in inducing the spring bloom. Lévy (2015) also used a 1D NPZD model to test different hypotheses and highlighted the importance of physics forcing on the validity of each hypothesis. However, few studies have considered the roles of the changes in phytoplankton traits and photoacclimation in determining the onset of spring bloom. Our model presents an exciting opportunity to fill this knowledge gap.

4.3.4 Understanding diel variations in phytoplankton cell properties

Another potential application of our model is to understand ~~the diel variations of~~ diel variations in cell size, abundance, and the cell cycle of phytoplankton in the oligotrophic ocean (Vaulot et al., 1995; Li et al., 2022). It is an interesting phenomenon that picophytoplankton cells such as *Prochlorococcus* and *Synechococcus* tend to show synchronized growth over a diel light/dark cycle. Phytoplankton cells tend to increase the size from sunrise to sunset but divide in late afternoon or evening, thus creating a mismatch between cell carbon production and abundance (Li et al., 2022). However, these diel rhythms can be different for different groups of phytoplankton (Vaulot and Marie, 1999). It is still unclear how these different patterns can be entirely explained by environmental (light, nutrient) variations or at least partly due to endogenous circadian clock (Heath and Spencer, 1985; Hellweger et al., 2020).

As ~~our model~~ PIBM is driven by a diel light/dark cycle, it can be used to understand what regulates the changes in cell properties linked to the cell cycle. Our model can be further modified to include cell cycles of phytoplankton to understand what controls phytoplankton division (Pascual and Caswell, 1997). Another promising future direction is to include more molecular processes such as gene expression and protein synthesis into the phytoplankton cell cycle, thus allowing us to link molecular studies with phytoplankton traits (Hellweger, 2020; Hellweger et al., 2020).

5 Conclusions

We introduce a novel 1D-hybrid Eulerian-Lagrangian model, uniquely tailored to explore how water column dynamics shape phytoplankton dynamics. Phytoplankton are ~~modelled-modeled~~ as super-individuals, a Lagrangian particle that represents a cluster of clonal phytoplankton cells that are physiologically identical and share a common history. Each phytoplankton super-individual is characterized by its cell size, temperature affinity and light affinity. Furthermore, these super-individuals possess the capability to mutate, enhancing the model's capacity to simulate phytoplankton growth, productivity, and diversity within dynamic aquatic environments.

The seasonal variability of temperature, irradiance, and vertical diffusivity at the BATS station enabled us to evaluate the response of our ecological model to environmental changes. By employing three master traits (size, temperature affinity, and light affinity), the individual-based model illustrates the evolution and ~~adaptation-acclimation~~ of the phytoplankton community to environmental conditions and the competition between different phytoplankton ecotypes. Furthermore, the model allows individual analysis, allowing us to scrutinize how each phytoplankton super-individual responds to the environmental conditions it encounters throughout its life cycle. ~~The model also~~ While the current version has several weaknesses, such as high computational costs, and the need for extensive parameterization, and validation. ~~However, with the appropriate experimental design,~~ it has several potential applications that would help us address questions related to the individual growth of phytoplankton, as well as the productivity and diversity of the phytoplankton community.

Code and data availability. The model code and input data are publicly available at <https://github.com/BingzhangChen/IBM>

under the MIT license and also available on zenodo <https://doi.org/10.5281/zenodo.13310586>.

1030 *Author contributions.* BC conceived the study and wrote the model code. IS wrote the first draft of the paper. Both authors contributed to a
later revision of the paper.

Competing interests. The authors declare that they have no conflict of interest.

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References

- Acevedo-Trejos, E., Brandt, G., Bruggeman, J., and Merico, A.: Mechanisms shaping size structure and functional diversity of phytoplankton communities in the ocean, *Sci. Rep.*, 5, 8918, <https://doi.org/10.1038/srep08918>, 2015.
- 1040 Acevedo-Trejos, E., Cadier, M., Chakraborty, S., Chen, B., Cheung, S. Y., Grigoratou, M., Guill, C., Hassenrück, C., Kerimoglu, O., Klausches, T., et al.: Modelling approaches for capturing plankton diversity (MODIV), their societal applications and data needs, *Front. Mar. Sci.*, 9, 975414, <https://doi.org/10.3389/fmars.2022.975414>, 2022.
- Ammerman, J. W., Hood, R. R., Case, D. A., and Cotner, J. B.: Phosphorus deficiency in the Atlantic: An emerging paradigm in oceanography, *EOS*, 84, 165–170, <https://doi.org/10.1029/2003EO180001>, 2003.
- 1045 Anderson, T., Gentleman, W., and Yool, A.: EMPOWER-1.0: an efficient model of planktonic ecosystems written in R, *Geosci. Model Dev.*, 8, 2231–2262, <https://doi.org/10.5194/gmd-8-2231-2015>, 2015.
- Baer, S. E., Lomas, M. W., Terpis, K. X., Mouginot, C., and Martiny, A. C.: Stoichiometry of *Prochlorococcus*, *Synechococcus*, and small eukaryotic populations in the western North Atlantic Ocean, *Environ. Microb.*, 19, 1568–1583, <https://doi.org/10.1111/1462-2920.13672>, 2017.
- Banas, N. S.: Adding complex trophic interactions to a size-spectral plankton model: Emergent diversity patterns and limits on predictability, *Ecol. Modell.*, 222, 2663–2675, <https://doi.org/10.1016/j.ecolmodel.2011.05.018>, 2011.
- 1050 Barkmann, W. and Woods, J.: On using a Lagrangian model to calibrate primary production determined from in vitro incubation measurements, *J. Plankton Res.*, <https://doi.org/10.1093/plankt/18.5.767>, 1996.
- Barton, A. D., Dutkiewicz, S., Flierl, G., Bragg, J., and Follows, M. J.: Patterns of diversity in marine phytoplankton, *Science*, 327, 1509–1511, <https://doi.org/10.1126/science.1184961>, 2010.
- 1055 Barton, S., Jenkins, J., Buckling, A., Schaum, C.-E., Smirnoff, N., Raven, J. A., and Yvon-Durocher, G.: Evolutionary temperature compensation of carbon fixation in marine phytoplankton, *Ecol. Lett.*, 23, 722–733, <https://doi.org/10.1111/ele.13469>, 2020.
- Baudry, J., Dumont, D., and Schloss, I. R.: Turbulent mixing and phytoplankton life history: a Lagrangian versus Eulerian model comparison, *Mar. Ecol. Prog. Ser.*, 600, 55–70, <https://doi.org/10.3354/meps12634>, 2018.
- Beardmore, R. E., Gudelj, I., Lipson, D. A., and Hurst, L. D.: Metabolic trade-offs and the maintenance of the fittest and the flattest, *Nature*, 1060 472, 342–346, <https://doi.org/10.1038/nature09905>, 2011.
- Beckmann, A., Schaum, C.-E., and Hense, I.: Phytoplankton adaptation in ecosystem models, *J. Theor. Biol.*, 468, 60–71, <https://doi.org/10.1016/j.jtbi.2019.01.041>, 2019.
- Behrenfeld, M. J.: Abandoning Sverdrup’s critical depth hypothesis on phytoplankton blooms, *Ecology*, 91, 977–989, <https://doi.org/10.1890/09-1207.1>, 2010.
- 1065 Behrenfeld, M. J., O’Malley, R. T., Boss, E. S., Westberry, T. K., Graff, J. R., Halsey, K. H., Milligan, A. J., Siegel, D. A., and Brown, M. B.: Revaluating ocean warming impacts on global phytoplankton, *Nat. Clim. Change*, 6, 323–330, <https://doi.org/10.1038/nclimate2838>, 2016.
- Bjedov, I., Tenaillon, O., Gérard, B., Souza, V., Denamur, E., Radman, M., Taddei, F., and Matic, I.: Stress-Induced Mutagenesis in Bacteria, *Science*, 300, 1404–1409, <https://doi.org/10.1126/science.1082240>, 2003.
- 1070 Blonder, B., Morrow, C. B., Maitner, B., Harris, D. J., Lamanna, C., Violle, C., Enquist, B. J., and Kerkhoff, A. J.: New approaches for delineating n-dimensional hypervolumes, *Methods Ecol. Evol.*, 9, 305–319, <https://doi.org/10.1111/2041-210X.12865>, 2018.

- Bolding, K. and Villarreal, M. R.: GOTM: A general ocean turbulence model: Theory, applications and test cases, Tech. rep., European Commission Tech. Rep. EUR 18745 EN, 1999.
- 1075 Brown, J. H., Gillooly, J. F., Allen, A. P., Savage, V. M., and West, G. B.: Toward a metabolic theory of ecology, *Ecology*, 85, 1771–1789, <https://doi.org/10.1890/03-9000>, 2004.
- Bruggeman, J. and Bolding, K.: A general framework for aquatic biogeochemical models, *Environ. Modell. Softw.*, 61, 249–265, <https://doi.org/10.1016/j.envsoft.2014.04.002>, 2014.
- Buitenhuis, E. T., Rivkin, R. B., Sailley, S., and Le Quéré, C.: Biogeochemical fluxes through microzooplankton, *Global Biogeochem. Cy.*, 24, <https://doi.org/10.1029/2009GB003601>, 2010.
- 1080 Carmona, C. P., de Bello, F., Mason, N. W., and Lepš, J.: Trait probability density (TPD): measuring functional diversity across scales based on TPD with R, *Ecology*, 100, e02 876, <https://doi.org/10.1002/ecy.2876>, 2019.
- Chen, B.: Thermal diversity affects community responses to warming, *Ecol. Modell.*, 464, 109 846, <https://doi.org/10.1016/j.ecolmodel.2021.109846>, 2022.
- Chen, B. and Laws, E. A.: Is there a difference of temperature sensitivity between marine phytoplankton and heterotrophs?, *Limnol. Oceanogr.*, 62, 806–817, <https://doi.org/10.1002/lno.10462>, 2017.
- 1085 Chen, B. and Liu, H.: Relationships between phytoplankton growth and cell size in surface oceans: Interactive effects of temperature, nutrients, and grazing, *Limnol. Oceanogr.*, 55, 965–972, <https://doi.org/10.4319/lo.2010.55.3.0965>, 2010.
- Chen, B. and Liu, H.: Comment: Unimodal relationship between phytoplankton-mass-specific growth rate and size: A reply to the comment by Sal and Lopez-Urrutia (2011), *Limnol. Oceanogr.*, 56, 1956–1958, <https://doi.org/10.4319/lo.2011.56.5.1956>, 2011.
- 1090 Chen, B., Wang, L., Song, S., Huang, B., Sun, J., and Liu, H.: Comparisons of picophytoplankton abundance, size, and fluorescence between summer and winter in northern South China Sea, *Cont. Shelf Res.*, 31, 1527–1540, <https://doi.org/10.1016/j.csr.2011.06.018>, 2011.
- Chen, B., Landry, M. R., Huang, B., and Liu, H.: Does warming enhance the effect of microzooplankton grazing on marine phytoplankton in the ocean?, *Limnology and oceanography*, 57, 519–526, 2012.
- Chen, B., Smith, S. L., and Wirtz, K. W.: Effect of phytoplankton size diversity on primary productivity in the North Pacific: trait distributions under environmental variability, *Ecol. Lett.*, 22, 56–66, <https://doi.org/10.1111/ele.13167>, 2019.
- 1095 Christensen, V. G., Olds, H. T., Norland, J., and Khan, E.: Phytoplankton community interactions and cyanotoxin mixtures in three recurring surface blooms within one lake, *J. Hazard. Mater.*, 427, 128 142, <https://doi.org/10.1016/j.jhazmat.2021.128142>, 2022.
- Cianelli, D., D’Alcalá, M. R., Saggiomo, V., and Zambianchi, E.: Coupling mixing and photophysiological response of Antarctic plankton: a Lagrangian approach, *Antarct. Sci.*, 16, 133–142, <https://doi.org/10.1017/s0954102004001968>, 2004.
- 1100 Cianelli, D., Sabia, L., d’Alcalà, M. R., and Zambianchi, E.: An individual-based analysis of the dynamics of two coexisting phytoplankton species in the mixed layer, *Ecol. Modell.*, 220, 2380–2392, <https://doi.org/10.1016/j.ecolmodel.2009.06.016>, 2009.
- Cloern, J. E.: Why large cells dominate estuarine phytoplankton, *Limnol. Oceanogr.*, 63, S392–S409, <https://doi.org/10.1002/lno.10749>, 2018.
- Cullen, J. J.: Subsurface chlorophyll maximum layers: enduring enigma or mystery solved?, *Annu. Rev. Mar. Sci.*, 7, 207–239, <https://doi.org/10.1146/annurev-marine-010213-135111>, 2015.
- 1105 De Bello, F., Carmona, C. P., Dias, A. T., Götzemberger, L., Moretti, M., and Berg, M. P.: Handbook of trait-based ecology: from theory to R tools, Cambridge University Press, 2021.
- Dell, A. I., Pawar, S., and Savage, V. M.: Systematic variation in the temperature dependence of physiological and ecological traits, *P. Natl. Acad. Sci. USA*, 108, 10 591–10 596, <https://doi.org/10.1073/pnas.1015178108>, 2011.

- 1110 Ducklow, H. W., Dickson, M.-L., Kirchman, D. L., Steward, G., Orchardo, J., Marra, J., and Azam, F.: Constraining bacterial production, conversion efficiency and respiration in the Ross Sea, Antarctica, January–February, 1997, *Deep-Sea Res. II*, 47, 3227–3247, [https://doi.org/10.1016/S0967-0645\(00\)00066-7](https://doi.org/10.1016/S0967-0645(00)00066-7), 2000.
- Durante, G., Basset, A., Stanca, E., and Roselli, L.: Allometric scaling and morphological variation in sinking rate of phytoplankton, *J. Phycol.*, 55, 1386–1393, <https://doi.org/10.1111/jpy.12916>, 2019.
- 1115 Dutkiewicz, S., Cermeño, P., Jahn, O., Follows, M. J., Hickman, A. E., Taniguchi, D. A., and Ward, B. A.: Dimensions of marine phytoplankton diversity, *Biogeosciences*, 17, 609–634, <https://doi.org/10.5194/bg-17-609-2020>, 2020.
- Dziekan, P. and Zmijewski, P.: University of Warsaw Lagrangian Cloud Model (UWLCM) 2.0: adaptation of a mixed Eulerian–Lagrangian numerical model for heterogeneous computing clusters, *Geosci. Model Dev.*, 15, 4489–4501, <https://doi.org/10.5194/gmd-15-4489-2022>, 2022.
- 1120 Edwards, K. F., Thomas, M. K., Klausmeier, C. A., and Litchman, E.: Allometric scaling and taxonomic variation in nutrient utilization traits and maximum growth rate of phytoplankton, *Limnol. Oceanogr.*, 57, 554–566, <https://doi.org/10.4319/lo.2012.57.2.0554>, 2012.
- Edwards, K. F., Thomas, M. K., Klausmeier, C. A., and Litchman, E.: Light and growth in marine phytoplankton: allometric, taxonomic, and environmental variation, *Limnol. Oceanogr.*, 60, 540–552, <https://doi.org/10.1002/lno.10033>, 2015.
- Falkowski, P. G. and Wirick, C. D.: A simulation model of the effects of vertical mixing on primary productivity, *Mar. Biol.*, 65, 69–75, <https://doi.org/10.1007/BF00397069>, 1981.
- 1125 Field, C. B., Behrenfeld, M. J., Randerson, J. T., and Falkowski, P.: Primary production of the biosphere: integrating terrestrial and oceanic components, *Science*, 281, 237–240, <https://doi.org/10.1126/science.281.5374.237>, 1998.
- Follows, M. J., Dutkiewicz, S., Grant, S., and Chisholm, S. W.: Emergent biogeography of microbial communities in a model ocean, *Science*, 315, 1843–1846, <https://doi.org/10.1126/science.1138544>, 2007.
- 1130 Gabriel, E., Fagg, G. E., Bosilca, G., Angskun, T., Dongarra, J. J., Squyres, J. M., Sahay, V., Kambadur, P., Barrett, B., Lumsdaine, A., Castain, R. H., Daniel, D. J., Graham, R. L., and Woodall, T. S.: Open MPI: Goals, Concept, and Design of a Next Generation MPI Implementation, in: *Proceedings, 11th European PVM/MPI Users’ Group Meeting*, pp. 97–104, Budapest, Hungary, https://doi.org/10.1007/978-3-540-30218-6_19, 2004.
- Gan, J., Lu, Z., Dai, M., Cheung, A. Y., Liu, H., and Harrison, P.: Biological response to intensified upwelling and to a river plume in the northeastern South China Sea: A modeling study, *J. Geophys. Res.-Oceans*, 115, <https://doi.org/10.1029/2009JC005569>, 2010.
- 1135 Garcia, H. E., Bouchard, C., Cross, S. L., Paver, C. R., Reagan, J. R., Boyer, T. P., Locarnini, R. A., Mishonov, A. V., Baranova, O. K., Seidov, D., et al.: *World Ocean Atlas 2023, Volume 4: Dissolved Inorganic Nutrients (Phosphate, Nitrate, and Silicate)*, NOAA Atlas Nesdis 92, p. 79pp, <https://doi.org/10.25923/39qw-7j08>, 2024.
- Gasol, J. M., Del Giorgio, P. A., and Duarte, C. M.: Biomass distribution in marine planktonic communities, *Limnol. Oceanogr.*, 42, 1353–1363, <https://doi.org/10.4319/lo.1997.42.6.1353>, 1997.
- 1140 Geider, R. J., MacIntyre, H. L., and Kana, T. M.: A dynamic regulatory model of phytoplanktonic acclimation to light, nutrients, and temperature, *Limnol. Oceanogr.*, 43, 679–694, <https://doi.org/10.4319/lo.1998.43.4.0679>, 1998.
- Grabowski, W. W., Morrison, H., Shima, S.-I., Abade, G. C., Dziekan, P., and Pawlowska, H.: Modeling of cloud microphysics: Can we do better?, *Bulletin of the American Meteorological Society*, 100, 655–672, <https://doi.org/10.1175/BAMS-D-18-0005.1>, 2019.
- 1145 Graff, J. R., Milligan, A. J., and Behrenfeld, M. J.: The measurement of phytoplankton biomass using flow-cytometric sorting and elemental analysis of carbon, *Limnol. Oceanogr.-Meth.*, 10, 910–920, <https://doi.org/10.4319/lom.2012.10.910>, 2012.

- Grimm, V., Berger, U., Bastiansen, F., Eliassen, S., Ginot, V., Giske, J., Goss-Custard, J. D., Grand, T. C., Heinz, S. K., Huse, G., Huth, A., Jepsen, J. U., Jorgensen, C., Mooij, W. M., Müller, B., Pe'er, G., Piou, C., Railsback, S. F., Robbins, A. M., Robbins, M. M., Rossmanith, E., Rüger, N., Strand, E., Souissi, S., Stillman, R. A., Vabø, R., Visser, U., and DeAngelis, D. L.: A standard protocol for describing individual-based and agent-based models, *Ecol. Modell.*, <https://doi.org/10.1016/j.ecolmodel.2006.04.023>, 2006.
- Grimm, V., Berger, U., DeAngelis, D. L., Polhill, G., Giske, J., and Railsback, S. F.: The ODD protocol: A review and first update, *Ecol. Modell.*, <https://doi.org/10.1016/j.ecolmodel.2010.08.019>, 2010.
- Han, B.-P.: A Mechanistic Model of Algal Photoinhibition Induced by Photodamage to Photosystem-II, *J. Theor. Biol.*, 214, 519–527, <https://doi.org/10.1006/jtbi.2001.2468>, 2002.
- Hansen, B., Bjørnsen, P. K., and Hansen, P. J.: The size ratio between planktonic predators and their prey, *Limnol. Oceanogr.*, 39, 395–403, <https://doi.org/10.4319/lo.1994.39.2.0395>, 1994.
- Hansen, P. J., Bjørnsen, P. K., and Hansen, B. W.: Zooplankton grazing and growth: Scaling within the 2–2,000- μm body size range, *Limnol. Oceanogr.*, 42, 687 – 704, <https://doi.org/10.4319/lo.1997.42.4.0687>, 1997.
- Harris, R., Wiebe, P., Lenz, J., Skjoldal, H.-R., and Huntley, M.: ICES Zooplankton Methodology Manual, Elsevier, ISBN 978-0-08-049533-0, 2000.
- Heath, M. and Spencer, C.: A model of the cell cycle and cell division phasing in a marine diatom, *Microbiology*, 131, 411–425, <https://doi.org/10.1099/00221287-131-3-411>, 1985.
- Hellweger, F. L.: Combining molecular observations and microbial ecosystem modeling: a practical guide, *Annu. Rev. Mar. Sci.*, 12, 267–289, <https://doi.org/10.1146/annurev-marine-010419-010829>, 2020.
- Hellweger, F. L., Kravchuk, E. S., Novotny, V., and Gladyshev, M. I.: Agent-based modeling of the complex life cycle of a cyanobacterium (*Anabaena*) in a shallow reservoir, *Limnol. Oceanogr.*, 53, 1227–1241, <https://doi.org/10.4319/lo.2008.53.4.1227>, 2008.
- Hellweger, F. L., van Sebille, E., and Fredrick, N. D.: Biogeographic patterns in ocean microbes emerge in a neutral agent-based model, *Science*, 345, 1346–1349, <https://doi.org/10.1126/science.1254421>, 2014.
- Hellweger, F. L., Jabbur, M. L., Johnson, C. H., van Sebille, E., and Sasaki, H.: Circadian clock helps cyanobacteria manage energy in coastal and high latitude ocean, *ISME J.*, 14, 560–568, <https://doi.org/10.1038/s41396-019-0547-0>, 2020.
- Hense, I. and Beckmann, A.: The representation of cyanobacteria life cycle processes in aquatic ecosystem models, *Ecol. Modell.*, 221, 2330–2338, <https://doi.org/10.1016/j.ecolmodel.2010.06.014>, 2010.
- Herrig, R. and Falkowski, P. G.: Nitrogen limitation in *Isochrysis galbana* (haptophyceae). i. photosynthetic energy conversion and growth efficiencies 1, *J. Phycol.*, 25, 462–471, <https://doi.org/10.1111/j.1529-8817.1989.tb00251.x>, 1989.
- Hooper, D. U., Chapin III, F. S., Ewel, J. J., Hector, A., Inchausti, P., Lavorel, S., Lawton, J. H., Lodge, D. M., Loreau, M., Naeem, S., et al.: Effects of biodiversity on ecosystem functioning: a consensus of current knowledge, *Ecological monographs*, 75, 3–35, 2005.
- Huete-Ortega, M., Rodríguez-Ramos, T., López-Sandoval, D. C., Cermeño, P., Blanco, J. M., Palomino, R. L., Rodríguez, J., and Marañón, E.: Distinct patterns in the size-scaling of abundance and metabolism in coastal and open-ocean phytoplankton communities, *Mar. Ecol. Prog. Ser.*, 515, 61–71, <https://doi.org/10.3354/meps11007>, 2014.
- Huston, M. A.: Biological diversity: the coexistence of species, Cambridge university press, 1994.
- Ivlev, V.: Experimental ecology of the feeding fishes, *Pishepromizdat*, 1955.
- Jokulsdottir, T. and Archer, D.: A stochastic, Lagrangian model of sinking biogenic aggregates in the ocean (SLAMS 1.0): model formulation, validation and sensitivity, *Geoscientific Model Development*, 9, 1455–1476, 2016.

Key, T., McCarthy, A., Campbell, D. A., Six, C., Roy, S., and Finkel, Z. V.: Cell size trade-offs govern light exploitation strategies in marine
1185 phytoplankton, *Environ. Microbiol.*, 12, 95–104, <https://doi.org/10.1111/j.1462-2920.2009.02046.x>, 2010.

Kida, S. and Ito, T.: A Lagrangian view of spring phytoplankton blooms, *Journal of Geophysical Research: Oceans*, 122, 9160–9175, 2017.

Kjørboe, T.: A mechanistic approach to plankton ecology, Princeton Univ. Press., ISBN 9780691134222, 2009.

Kwiatkowski, L., Yool, A., Allen, J., Anderson, T., Barciela, R., Buitenhuis, E., Butenschön, M., Enright, C., Halloran, P., Le Quéré, C.,
1190 et al.: iMarNet: an ocean biogeochemistry model intercomparison project within a common physical ocean modelling framework, *Bio-
geosciences*, 11, 7291–7304, 2014.

Le Gland, G., Vallina, S. M., Smith, S. L., and Cermeño, P.: SPEAD 1.0-Simulating Plankton Evolution with Adaptive Dynamics in a two-trait
continuous fitness landscape applied to the Sargasso Sea, *Geosci. Model Dev.*, 14, 1949–1985, <https://doi.org/10.5194/gmd-14-1949-2021>,
2021.

Ledbetter, M.: Langmuir circulations and plankton patchiness, *Ecol. Modell.*, 7, 289–310, [https://doi.org/https://doi.org/10.1016/0304-
1195 3800\(79\)90039-5](https://doi.org/https://doi.org/10.1016/0304-3800(79)90039-5), 1979.

Lévy, M.: Exploration of the critical depth hypothesis with a simple NPZ model, *ICES J. Mar. Sci.*, 72, 1916–1925,
<https://doi.org/10.1093/icesjms/fsv016>, 2015.

Li, C., Chiang, K.-P., Laws, E. A., Liu, X., Chen, J., Huang, Y., Chen, B., Tsai, A.-Y., and Huang, B.: Quasi-Antiphase Diel Patterns
of Abundance and Cell Size/Biomass of Picophytoplankton in the Oligotrophic Ocean, *Geophys. Res. Lett.*, 49, e2022GL097753,
1200 <https://doi.org/10.1029/2022GL097753>, 2022.

Litchman, E., Klausmeier, C. A., and Yoshiyama, K.: Contrasting size evolution in marine and freshwater diatoms, *P. Natl. Acad. Sci. USA*,
106, 2665–2670, <https://doi.org/10.1073/pnas.0810891106>, 2009.

Locarnini, R. A., Mishonov, A. V., Baranova, O. K., Reagan, J. R., Boyer, T. P., Seidov, D., Wang, Z., Garcia, H. E., Bouchard, C., Cross,
S. L., et al.: World Ocean Atlas 2023, Volume 1: Temperature, NOAA Atlas NESDIS 89, p. 52pp, <https://doi.org/10.25923/54bh-1613>,
1205 2024.

Marañón, E.: Phytoplankton Size Structure, in: *Encyclopedia of Ocean Sciences (Third Edition)*, edited by Cochran, J. K., Bokuniewicz,
H. J., and Yager, P. L., pp. 599–605, Academic Press, Oxford, third edition edn., ISBN 978-0-12-813082-7, [https://doi.org/10.1016/B978-
0-12-409548-9.11405-8](https://doi.org/10.1016/B978-0-12-409548-9.11405-8), 2019.

Marañón, E., Cermeño, P., López-Sandoval, D. C., Rodríguez-Ramos, T., Sobrino, C., Huete-Ortega, M., Blanco, J. M., and Rodríguez,
1210 J.: Unimodal size scaling of phytoplankton growth and the size dependence of nutrient uptake and use, *Ecol. Lett.*, 16, 371–379,
<https://doi.org/10.1111/ele.12052>, 2013.

Mason, N. W., Mouillot, D., Lee, W. G., and Wilson, J. B.: Functional richness, functional evenness and functional divergence: the primary
components of functional diversity, *Oikos*, 111, 112–118, <https://doi.org/10.1111/j.0030-1299.2005.13886.x>, 2005.

Mayzaud, P. and Poulet, S. A.: The importance of the time factor in the response of zooplankton to varying concentrations of naturally
1215 occurring particulate matter, *Limnol. Oceanogr.*, 23, 1144–1154, <https://doi.org/10.4319/lo.1978.23.6.1144>, 1978.

Menden-Deuer, S. and Lessard, E. J.: Carbon to volume relationships for dinoflagellates, diatoms, and other protist plankton, *Limnol.
Oceanogr.*, 45, 569–579, <https://doi.org/10.4319/lo.2000.45.3.0569>, 2000.

Moore, L. R., Rocap, G., and Chisholm, S. W.: Physiology and molecular phylogeny of coexisting *Prochlorococcus* ecotypes, *Nature*, 393,
464–467, <https://doi.org/10.1038/30965>, 1998.

Morel, F. M.: Kinetics of nutrient uptake and growth in phytoplankton, *J. Phycol.*, 23, <https://doi.org/10.1111/j.0022-3646.1987.00137.x>,
1220 1987.

- Nikolaou, A., Hartmann, P., Sciandra, A., Chachuat, B., and Bernard, O.: Dynamic coupling of photoacclimation and photoinhibition in a model of microalgae growth, *J. Theor. Biol.*, 390, 61–72, <https://doi.org/10.1016/j.jtbi.2015.11.004>, 2016.
- 1225 Nogueira, E., Woods, J., Harris, C., Field, A., and Talbot, S.: Phytoplankton co-existence: Results from an individual-based simulation model, *Ecol. Modell.*, 198, 1–22, <https://doi.org/10.1016/j.ecolmodel.2006.04.013>, 2006.
- Noh, K. M., Noh, Y., Brereton, A., and Kug, J.-S.: The route to spring phytoplankton blooms simulated by a Lagrangian plankton model, *Journal of Geophysical Research: Oceans*, 126, e2020JC016 753, 2021.
- Norberg, J., Swaney, D. P., Dushoff, J., Lin, J., Casagrandi, R., and Levin, S. A.: Phenotypic diversity and ecosystem functioning in changing environments: a theoretical framework, *P. Natl. Acad. Sci. USA*, 98, 11 376–11 381, <https://doi.org/10.1073/pnas.171315998>, 2001.
- 1230 Pahlow, M., Dietze, H., and Oschlies, A.: Optimality-based model of phytoplankton growth and diazotrophy, *Mar. Ecol. Prog. Ser.*, 489, 1–16, <https://doi.org/10.3354/meps10449>, 2013.
- Pascual, M. and Caswell, H.: From the cell cycle to population cycles in phytoplankton–nutrient interactions, *Ecology*, 78, 897–912, <https://doi.org/10.2307/2266068>, 1997.
- Platt, T. and Denman, K.: Organisation in the pelagic ecosystem, *Helgoland Wiss. Meer.*, 30, 575–581, <https://doi.org/10.1007/BF02207862>,
 1235 1977.
- Platt, T. and Gallegos, C.: MOD PRIM PROD, in: *Primary Production in the Sea*, edited by Falkowski, P., pp. 339–362, Plenum, ISBN 978-1-4684-3890-1, <https://doi.org/10.1007/978-1-4684-3890-1>, 1981.
- Ranjbar, M. H., Hamilton, D. P., Etemad-Shahidi, A., and Helfer, F.: Individual-based modelling of cyanobacteria blooms: Physical and physiological processes, *Sci. Total Environ.*, 792, 148 418, <https://doi.org/10.1016/j.scitotenv.2021.148418>, 2021.
- 1240 Rao, C. R.: Diversity and dissimilarity coefficients: a unified approach, *Theoretical population biology*, 21, 24–43, 1982.
- Righetti, D., Vogt, M., Gruber, N., Psomas, A., and Zimmermann, N. E.: Global pattern of phytoplankton diversity driven by temperature and environmental variability, *Sci. Adv.*, 5, eaau6253, <https://doi.org/10.1126/sciadv.aau625>, 2019.
- Ross, O. and Geider, R.: New cell-based model of photosynthesis and photo-acclimation: accumulation and mobilisation of energy reserves in phytoplankton, *Mar. Ecol. Prog. Ser.*, 383, 53–71, <https://doi.org/10.3354/meps07961>, 2009.
- 1245 Ross, O. N. and Sharples, J.: Recipe for 1-D Lagrangian particle tracking models in space-varying diffusivity, *Limnol. Oceanogr.-Meth.*, 2, 289–302, <https://doi.org/10.4319/lom.2004.2.289>, 2004.
- Ross, O. N., Geider, R. J., Berdalet, E., Artigas, M. L., and Piera, J.: Modelling the effect of vertical mixing on bottle incubations for determining in situ phytoplankton dynamics. I. Growth rates, *Marine Ecology Progress Series*, 435, 13–31, 2011.
- Scheffer, M., Baveco, J., DeAngelis, D., Rose, K., and van Ness, E.: Super-individuals a simple solution for modelling large populations on
 1250 an individual basis, *Ecol. Modell.*, 80, 161 – 170, [https://doi.org/10.1016/0304-3800\(94\)00055-M](https://doi.org/10.1016/0304-3800(94)00055-M), 1995.
- Shchepetkin, A. F. and McWilliams, J. C.: The regional oceanic modeling system (ROMS): a split-explicit, free-surface, topography-following-coordinate oceanic model, *Ocean Modell.*, 9, 347–404, <https://doi.org/10.1016/j.ocemod.2004.08.002>, 2005.
- Sigman, D. M. and Boyle, E. A.: Glacial/interglacial variations in atmospheric carbon dioxide, *Nature*, 407, 859–869, <https://doi.org/10.1038/35038000>, 2000.
- 1255 Thomas, M. K., Kremer, C. T., Klausmeier, C. A., and Litchman, E.: A global pattern of thermal adaptation in marine phytoplankton, *Science*, 338, 1085–1088, <https://doi.org/10.1126/science.1224836>, 2012.
- Tomkins, M., Martin, A. P., Nurser, A. G., and Anderson, T. R.: Phytoplankton acclimation to changing light intensity in a turbulent mixed layer: A Lagrangian modelling study, *Ecol. Modell.*, 417, 108 917, <https://doi.org/10.1016/j.ecolmodel.2019.108917>, 2020.

- Trebilco, R., Baum, J. K., Salomon, A. K., and Dulvy, N. K.: Ecosystem ecology: size-based constraints on the pyramids of life, *Trends Ecol. Evol.*, 28, 423–431, <https://doi.org/10.1016/j.tree.2013.03.008>, 2013.
- Vallina, S. M., Follows, M. J., Dutkiewicz, S., Montoya, J. M., Cermeño, P., and Loreau, M.: Global relationship between phytoplankton diversity and productivity in the ocean, *Nat. Commun.*, 5, 4299, <https://doi.org/10.1038/ncomms5299>, 2014.
- Vaulot, D. and Marie, D.: Diel variability of photosynthetic picoplankton in the equatorial Pacific, *J. Geophys. Res.-Oceans*, 104, 3297–3310, <https://doi.org/10.1029/98JC01333>, 1999.
- Vaulot, D., Marie, D., Olson, R. J., and Chisholm, S. W.: Growth of *Prochlorococcus*, a photosynthetic prokaryote, in the equatorial Pacific Ocean, *Science*, 268, 1480–1482, <https://doi.org/10.1126/science.268.5216.1480>, 1995.
- Vellend, M. and Agrawal, A.: Conceptual synthesis in community ecology, *Q. Rev. Biol.*, 85, 183–206, <https://doi.org/10.1086/652373>, 2010.
- Visser, A. W.: Using random walk models to simulate the vertical distribution of particles in a turbulent water column, *Mar. Ecol. Prog. Ser.*, 158, 275–281, <https://doi.org/10.3354/meps158275>, 1997.
- Ward, B. A., Dutkiewicz, S., Jahn, O., and Follows, M. J.: A size-structured food-web model for the global ocean, *Limnol. Oceanogr.*, 57, 1877–1891, <https://doi.org/10.4319/lo.2012.57.6.1877>, 2012.
- Ward, B. A., Collins, S., Dutkiewicz, S., Gibbs, S., Bown, P., Ridgwell, A., Sauterey, B., Wilson, J., and Oschlies, A.: Considering the role of adaptive evolution in models of the ocean and climate system, *J. Adv. Model Earth. Sy.*, 11, 3343–3361, <https://doi.org/10.1029/2018MS001452>, 2019.
- Wirtz, K. W.: Non-uniform scaling in phytoplankton growth rate due to intracellular light and CO₂ decline, *J. Plankton. Res.*, 33, 1325–1341, <https://doi.org/10.1093/plankt/fbr021>, 2011.
- Woods, J.: The Lagrangian Ensemble metamodel for simulating plankton ecosystems, *Prog. Oceanogr.*, <https://doi.org/10.1016/j.pocean.2005.04.003>, 2005.
- Woods, J. and Barkmann, W.: Diatom demography in winter – simulated by the Lagrangian Ensemble method, *Fish. Oceanogr.*, 2, 202–222, <https://doi.org/10.1111/j.1365-2419.1993.tb00136.x>, 1993.
- Woods, J. and Onken, R.: Diurnal variation and primary production in the ocean preliminary results of a Lagrangian ensemble model, *J. Plankton. Res.*, 4, 735–756, <https://doi.org/10.1093/plankt/4.3.735>, 1982.
- Woods, J. D. and Barkmann, W.: Simulating plankton ecosystems by the Lagrangian Ensemble Method, *Philos. T. Roy. Soc. B*, <https://doi.org/10.1098/rstb.1994.0004>, 1994.
- Wu, Z. and Forget, G.: PlanktonIndividuals.jl: a GPU supported individual-based phytoplankton life cycle model, *Journal of Open Source Software*, 7, 4207, 2022.