



1 **A New Simplified Parameterization of Secondary Organic Aerosol in**
2 **the Community Earth System Model Version 2 (CESM2; CAM6.3)**

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9 **Abstract.** The Community Earth System Model (CESM) community has been providing versatile
10 modeling options, with simple to complex chemistry and aerosol schemes in a single model, in order to
11 support the broad scientific community with various research interests. While different model
12 configurations are available in CESM and these can be used for different fields of Earth system science,
13 simulation results that are consistent across configurations are still desirable. Here we develop a new
14 simple secondary organic aerosol (SOA) scheme in the Community Atmosphere Model version 6.3
15 (CAM6.3), the atmospheric component of the CESM. The main purpose of this simplified SOA scheme
16 is to reduce the differences in aerosol concentrations and radiative fluxes between CAM and CAM with
17 detailed chemistry (CAM-chem) while maintaining the computational efficiency of CAM. CAM
18 simulation results with the current and the new SOA schemes are compared to the CAM-chem results as
19 a reference. More consistent SOA concentrations are obtained globally when using the new SOA
20 scheme, for both temporal and spatial variabilities. Furthermore, the overestimation of other
21 carbonaceous aerosols (black carbon and primary organic aerosol) in CAM is greatly reduced, which
22 results in improved global atmospheric burden and concentrations at the high latitudes of the Northern
23 Hemisphere compared to the full chemistry version (CAM-chem). As a result, the high bias of radiative
24 flux in the Arctic region is significantly reduced for both nudged and free-running simulations. We find
25 that the current SOA scheme in CAM can still be used for radiative forcing calculation as the high
26 biases exist both in pre-industrial and present conditions, but studies focusing on the instantaneous
27 radiative effects would benefit from using the new SOA scheme. The new SOA scheme also has
28 technical advantages including the use of identical SOA precursor emissions as CAM-chem from the
29 online biogenic emissions, instead of pre-calculated emissions that may introduce differences. Future
30 parameter updates on the CAM-chem SOA scheme can be easily translated to the new CAM SOA
31 scheme as it is derived from the CAM-chem SOA scheme.



32 **Short Summary.** The new simple secondary organic aerosol (SOA) scheme has been developed for the
33 Community Atmosphere Model (CAM), based on the complex SOA scheme in CAM with detailed
34 chemistry (CAM-chem). The CAM with the new SOA scheme shows better agreements with
35 CAM-chem in terms of aerosol concentrations and radiative fluxes, which ensures more consistent
36 results between different compsets in the Community Earth System Model. The new SOA scheme also
37 has technical advantages for future developments.



38 **1 Introduction**

39 Secondary organic aerosol (SOA) accounts for a substantial fraction of ambient tropospheric
40 aerosol (Hallquist et al., 2009). Atmospheric models generally use parameterizations to simulate SOA
41 because it is composed of a wide range of different organic molecules (Goldstein and Galbally, 2007).
42 The SOA parameterization in atmospheric chemistry models varies from the simple method of
43 multiplying constant yields to emissions, to the complex volatility basis set (VBS) approach, which
44 considers the oxidation of volatile organic compounds (VOCs) and gas-particle partitioning, as shown
45 in the recent model intercomparison study for organic aerosol (OA) (Hodzic et al., 2020).

46 Climate models that have to perform hundreds of years of simulations and many ensemble
47 members often use very simple parameterizations to calculate SOA in the model (Tsigaridis and
48 Kanakidou, 2018), due to the high computational cost associated with chemistry, deposition, and the
49 increased number of model tracers to be transported (Jo et al., 2019). Because SOA affects climate
50 through aerosol-radiation and aerosol-cloud interactions, and climate also affects SOA through
51 changing biogenic emissions and photochemistry (Gottelman et al., 2019a; Sporre et al., 2019; Tilmes et
52 al., 2019; Jo et al., 2021), the accurate representation of SOA in climate models is important but needs
53 to have low computational cost for long-term simulation purposes.

54 The Community Earth System Model Version 2 (CESM2) has two different SOA schemes, one
55 simplified scheme for the Community Atmosphere Model version 6 (CAM6) (Danabasoglu et al., 2020)
56 and the Whole Atmosphere Community Climate Model version 6 (WACCM6) with the Middle
57 Atmosphere (MA) chemistry (Gottelman et al., 2019b), and a VBS scheme for the CAM6 with
58 comprehensive chemistry (CAM6-chem) (Emmons et al., 2020) and the WACCM6 with the TSMLT
59 (troposphere, stratosphere, mesosphere, and lower thermosphere) mechanism. For the purpose of
60 climate studies using many ensemble members, CAM6 is generally used for computational efficiency.
61 Models like WACCM6 with TSMLT are used for detailed chemistry and aerosol studies, but in general,
62 only a few ensemble members can be performed. Ideally, the two SOA schemes in simple and complex
63 chemistry configurations should give the same results to maintain model consistency regarding aerosol
64 fields and resulting climate forcings, but the spatial and temporal distributions of SOA between CAM



65 and CAM-chem (and WACCM TSMLT) are different enough to have a significant effect on black
 66 carbon (BC) and the Earth’s radiation budget (Tilmes et al., 2019).

67 Here we propose a new simplified and computationally affordable SOA scheme for CAM, which is
 68 based on the VBS scheme in CAM-chem. We compare three SOA schemes (VBS, current simplified
 69 SOA, and new simplified SOA) under a few different CESM2 configurations (specified dynamics and
 70 free-running in preindustrial and present conditions). The new approach substantially reduces the
 71 differences in aerosol and radiation values between CAM and CAM-chem (Sect. 3). The new SOA
 72 scheme also has a technical advantage as it does not need input files for the SOA precursor, but uses the
 73 same emissions files as CAM-chem or WACCM for individual SOA precursor species (isoprene,
 74 terpenes, toluene, etc.).

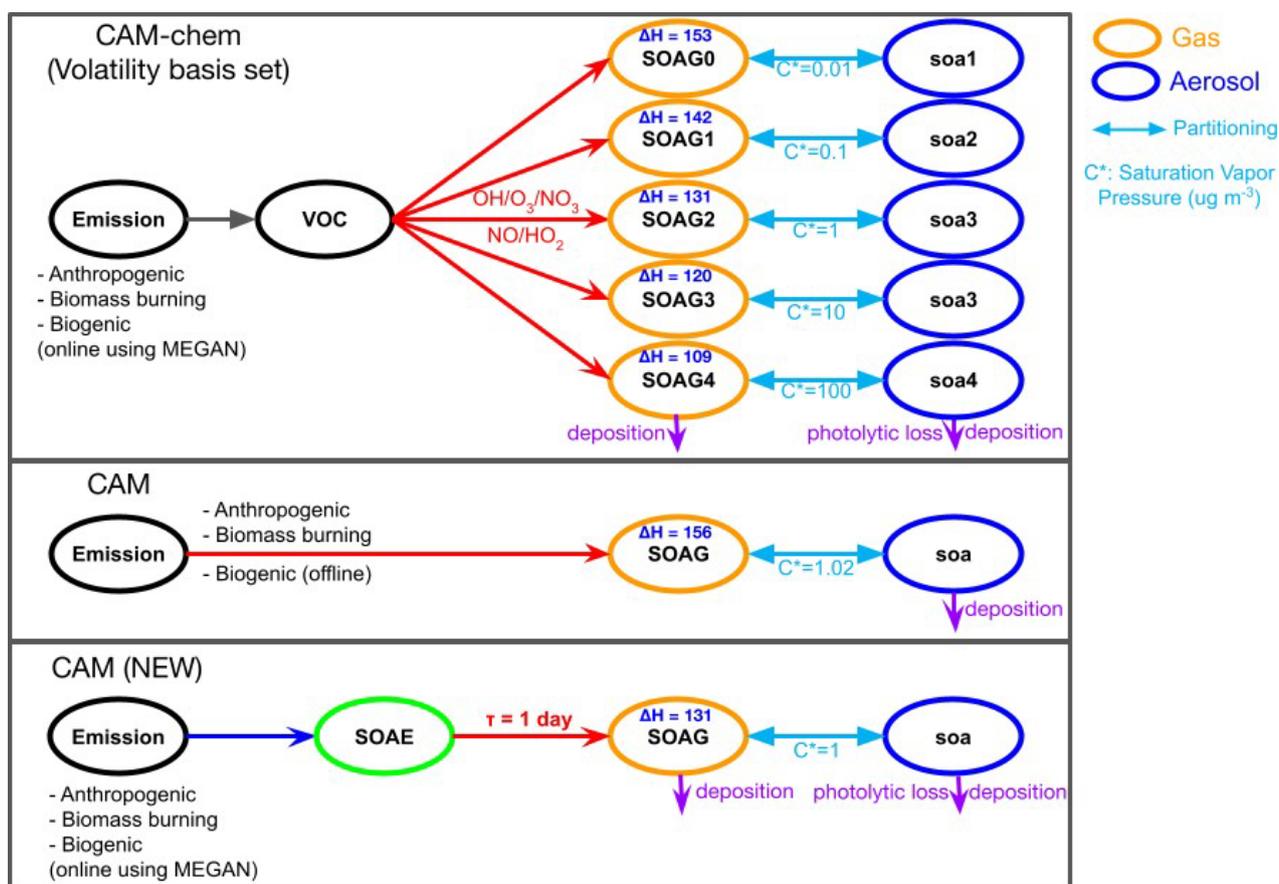
75 **Table 1.** SOA schemes used in this study. Computational costs are estimated on the Cheyenne
 76 supercomputer at NCAR. Computational cost ranges are given in parentheses with the average value.

SOA scheme	CAM-chem	CAM	CAM (NEW)
Emissions	Individual VOCs, online biogenic emissions	Pre-calculated, lumped SOAG emissions	Individual VOCs, online biogenic emissions
VOCs and chemistry	explicitly simulated	No	Lumped tracer (SOAE) with 1-day lifetime
Number of SOA bins	5	1	1
Saturation vapor pressure	0.01, 0.1, 1, 10, 100	1.02	1
Enthalpy of vaporization	153, 142, 131, 120, 109	156	131
SOA yield	Based on the VBS	Fixed fraction and scaled up by 50%	Based on the VBS but lumped
Loss processes	wet & dry deposition of SOAG photolytic loss of SOA	No deposition of SOAG No photolytic loss	wet & dry deposition of SOAG photolytic loss of SOA
Computational cost (pe-hrs / simulated_year)	7933 (7783 - 8083)	2398 (2353 - 2448)	2455 (2414 - 2501)



93 **2 Method**

94 In this section, we present SOA schemes in CAM-chem and CAM, along with the new simplified
 95 SOA scheme, as summarized in Fig. 1 and Table 1. General descriptions for other carbonaceous
 96 aerosols (BC and primary organic aerosol (POA)) are also explained as concentrations of those
 97 carbonaceous aerosols are affected by SOA concentrations (Tilmes et al., 2019). This section also
 98 includes the simulation set-up for comparisons between SOA schemes in Sect 3.



99 **Figure 1.** Schematic diagrams of SOA parameterizations in CESM2.



100 2.1 SOA scheme in CAM-chem

101 SOA in CAM-chem is simulated using the VBS approach, as described by Tilmes et al. (2019). The
102 VBS scheme in CAM-chem incorporates recent findings such as wall-corrected SOA yields, photolytic
103 removal of SOA, and more efficient removal by dry and wet deposition. Details can be found in Hodzic
104 et al. (2016). Here we briefly describe the characteristics that can be compared to the simple SOA
105 scheme in CAM.

106 CAM-chem uses a VBS scheme with 5 volatility bins (see Fig. 1) with saturation vapor pressures
107 spanning from 0.01 to 100 $\mu\text{g m}^{-3}$ at 300K. Enthalpy of vaporization values are 153, 142, 131, 120, and
108 109 kJ mol^{-1} for 0.01, 0.1, 1, 10, and 100 $\mu\text{g m}^{-3}$, respectively, at 300K based on Epstein et al. (2010).
109 Traditional SOA precursors such as isoprene, monoterpenes, sesquiterpenes, benzene, toluene, and
110 xylenes are explicitly simulated in the model, and the oxidation of those VOCs with OH, O₃, and NO₃
111 makes gas phase intermediate precursors of SOA (SOAG) according to the volatility bins. Semi- and
112 intermediate-range volatility organic compounds (S/IVOCs) are also considered with a simple OH
113 reaction. Since S/IVOCs are defined by volatility and exact chemical speciation is not available for
114 them, 60% of POA and 20% of total non-methane VOC (NMVOC) emissions are assumed to be
115 SVOCs and IVOCs, respectively (Hodzic et al., 2016). Biogenic VOCs are calculated online using the
116 model of emissions of gases and aerosols from nature version 2.1 (MEGAN2.1) (Guenther et al., 2012)
117 available in the Community Land Model version 5 (CLM5), a component of CESM and coupled to
118 CAM (Lawrence et al., 2019). CAM-chem also supports an extended VBS compset that keeps track of
119 VBS tracers from three sources (anthropogenic, biomass burning, and biogenic), leading to 15 SOA
120 species simulated in total. This option is not generally used except for studies tracking sources of SOA,
121 as total SOA burden and formation are very similar between the two options because the same volatility
122 bins are used (Tilmes et al., 2019).

123 In terms of aerosol modes, the four-mode version of the Modal Aerosol Module (MAM4) is
124 generally used in recent scientific applications (Liu et al., 2016). MAM4 considers Aitken,
125 accumulation, coarse, and primary carbon modes. SOA is simulated using Aitken and accumulation
126 modes but most of the mass (>99%) is in the accumulation mode (Tilmes et al., 2019). In total, 15



127 tracers (5 for the gas phase and 10 for the aerosol phase - 5 bins \times 2 modes) are used for the SOA
128 calculation in CAM-chem.

129 **2.2 Current SOA scheme in CAM**

130 The current simplified SOA scheme in CAM uses 3 tracers (1 for the gas phase and 2 for the
131 aerosol phase). Like the VBS, both gas-phase (SOAG) and aerosol-phase (soa_a1 and soa_a2 for
132 accumulation and Aitken modes) are simulated with gas-aerosol partitioning, with the enthalpy of
133 vaporization of 156 kJ mol⁻¹ and the saturation vapor pressure of 1.02 $\mu\text{g m}^{-3}$ (Liu et al., 2012). SOAG
134 does not undergo dry and wet removal, which is also different from the VBS that calculates dry and wet
135 deposition of gas-phase semivolatiles.

136 Unlike the VBS representation which explicitly simulates parent VOCs, this scheme does not
137 simulate the chemistry of VOCs but uses pre-calculated emissions using fixed mass yields for the
138 following VOC categories: 5% BIGALK (lumped $\geq\text{C}_4$ alkanes), 5% BIGENE (lumped $\geq\text{C}_4$ alkenes),
139 15% aromatics, 4% isoprene, and 25% monoterpenes (Liu et al., 2012). For biogenic VOCs, offline
140 emissions are precalculated and provided as an additional input file based on biogenic emissions
141 simulated by CLM-MEGAN2.1. Generally, the offline biogenic VOC emission does not have annual
142 variations and is repeated over the simulation period. Note that those SOAG emissions are further
143 increased by 50% after model tuning involving aerosol indirect effect (Liu et al., 2012).

144 **2.3 New SOA scheme in CAM**

145 The new SOA scheme uses a similar approach to the current SOA scheme in CAM, but several
146 modifications have been made to allow more consistent results with the VBS scheme in CAM-chem.
147 First, VOC species that generate SOA are matched to the VBS. In other words, BIGALK and BIGENE
148 are no longer used for the calculation of SOA emissions, and instead, sesquiterpenes and S/IVOCs are
149 considered for calculating the interactive emissions of SOA. This change can be scientifically justified
150 because SOA yields increase with the carbon number (Srivastava et al., 2022). BIGALK and BIGENE
151 are mainly composed of C4-C6 alkanes and alkenes (Emmons et al., 2020), but S/IVOCs correspond to
152 C12 or higher n-alkanes (Robinson et al., 2007).



153 Second, SOA yields (used for the interactive emissions) have been calculated based on the VBS
154 yields in CAM-chem. SOA yields for the first four bins and 20% of the fifth bin are summed up for
155 each compound. Only 20% of the fifth bin yield is used, as it is the most volatile bin. We consider SOA
156 yields from OH reactions only in this calculation, because reaction with OH is dominant for VOCs.
157 Only low NO_x yields are used in this study which is consistent with Tilmes et al. (2019), which is
158 appropriate for global climate studies with 1° horizontal resolution of the model grid. For air quality
159 studies with high spatial resolution, CAM-chem with NO_x-dependent SOA yields can be used
160 (Schwantes et al., 2022). The resulting SOA yields derived from CAM-chem results are 0.28, 0.64,
161 0.04, 0.16, 0.45, 0.35, 0.41, and 0.80 for monoterpenes, sesquiterpenes, isoprene, benzene, toluene,
162 xylenes, IVOC, and SVOC, respectively. Those are constants and do not change during the run, like
163 SOA yields used in the VBS scheme. It is worth noting that those yields can be easily updated in the
164 CAM run-time namelist file if there is a future update to the CAM-chem VBS scheme.

165 Third, we add a new tracer called “SOAE” (Fig. 1) to consider the time that VOCs and intermediate
166 chemical species undergo oxidation before forming semivolatiles. We assume a constant 1-day e-folding
167 lifetime to convert “SOAE” to “SOAG” which can be partitioned into aerosols so that oxidant fields do
168 not have to be simulated in CAM for computational efficiency. The 1-day lifetime corresponds to the
169 OH reaction rate constant of 10¹² cm³ molecules⁻¹ s⁻¹ with a global annual mean OH concentration of
170 11.6 × 10⁵ molecules cm⁻³ (Warneck and Williams, 2012).

171 Fourth, parameters are adjusted for consistency with the VBS scheme. The enthalpy of vaporization
172 is changed from 156 to 131 kJ mol⁻¹, which is the value used in the third bin of the VBS scheme. This
173 can change SOA in the upper troposphere where temperature dependency becomes important.
174 Deposition of gas phase SOA (SOAG) and the photolytic reaction of SOA are also added, which can
175 affect SOA concentrations in the remote atmosphere. Saturation vapor pressure change with the
176 assumption of 10% of POA as oxygenated (Liu et al., 2012) is not used in this scheme for consistency
177 with the VBS scheme.

178 Fifth, the same offline emission files (anthropogenic and biomass burning) and online emission
179 (biogenic) are used as the VBS method, via namelist control. As a result, preprocessing for SOAG
180 emission is no longer needed, and annual variability as well as the diurnal cycle for biogenic emission



181 can be easily considered. Note that biogenic emission is always calculated in CLM, regardless of
182 whether the emission is used or not in CAM or CAM-chem. Therefore, using online biogenic emissions
183 does not add computational cost.

184 **2.4 Other carbonaceous aerosols**

185 Here we describe BC and POA simulations in CAM and CAM-chem, as those are affected by SOA
186 concentrations through microphysics. Because BC, POA, and SOA precursors are emitted from the
187 same sources (except for the biogenic SOA) and share the same aerosol mode (accumulation mode),
188 changes in one component can significantly affect other components. Tilmes et al. (2019) reported
189 ~20% differences between the simplified SOA and the VBS scheme in terms of the global burden of BC
190 and POA, while the difference for the sulfate burden was very small (< 1%).

191 Unlike SOA, there is no difference in BC and POA simulation schemes between CAM and
192 CAM-chem, because BC and POA are chemically inert and the standard aerosol module is the same
193 (MAM4) for both CAM and CAM-chem. BC and POA are emitted into the primary carbon mode
194 (bc_a4 and pom_a4) and are transferred to the accumulation mode (bc_a1 and pom_a1) through
195 microphysical aging (Liu et al., 2016). The aging rate is substantially affected by SOA, through
196 changing internally mixed aerosol numbers (Tilmes et al., 2019).

197 **2.5 Simulation set-up**

198 We conduct three types of model experiments for different application scenarios using the
199 development version of CESM2.2 or CAM6.3 (tag name: cam6_3_050). First, a specified dynamics run
200 is performed for the analysis of the present condition using the nudged meteorological fields.
201 Temperature and horizontal winds are nudged towards the Modern-Era Retrospective analysis for
202 Research and Applications version 2 (MERRA2) every 3 hours (Gelaro et al., 2017). In this simulation,
203 we run the model for the year 2013 with a spin-up period of one year. Second, historical runs are
204 performed for the 1850s and 2000s with prescribed sea surface temperatures and sea ice conditions.
205 These are free-running simulations for 12 years for each condition with the two years discarded for the
206 spin-up. In this case, the CLM is run with the satellite phenology (SP) option which uses a prescribed



207 leaf area index (LAI) based on MODIS satellite observations (Lawrence et al., 2019). In this option, the
208 input LAI value for each plant functional type (PFT) is the same between the 1850s and 2000s but the
209 PFT fraction changes with time. As a result, the final LAI used for biogenic emission calculation is
210 slightly different between the two periods. The third is the same as the second experiment, but the
211 vegetation state including LAI is simulated prognostically by CLM (biogeochemistry; BGC) (Lawrence
212 et al., 2019). In addition to absolute values, the difference between the 1850s and 2000s is investigated
213 from the historical simulations in Sect. 3.3, to compare simulation results in terms of the radiative
214 forcing. In all simulations, the bi-directional oceanic flux of dimethyl sulfide (DMS) is calculated using
215 the Online Air-Sea Interface for Soluble Species (OASISS) (Wang et al., 2019, 2020) and the
216 climatological surface seawater DMS concentration (Lana et al., 2011), which will be the default DMS
217 emission in the next CESM version (CAM7).

218 **3 Results**

219 In this section, the current and new SOA schemes in CAM are evaluated against CAM-chem as a
220 reference. As SOA changes can affect the other carbonaceous aerosols and radiation fields in CESM2
221 (Tilmes et al., 2019), we also compare those simulation fields as shown in Table 2.

222 **3.1 Aerosols**

223 Table 2 shows the global annual mean burden of aerosols by different simulations, including
224 gas-phase SOA or semivolatiles (SOAG). Two CAM cases and CAM-chem are consistent within 10%
225 in terms of global SOA burden, with the new scheme showing better agreement. gas-phase SOA
226 (SOAG) is substantially underestimated in both CAM cases, because high volatility bins (saturation
227 vapor pressure of $10 \mu\text{g m}^{-3}$ and $100 \mu\text{g m}^{-3}$) are not simulated in the 1-bin simple SOA scheme.
228 However, SOAG does not affect other aerosol concentrations and radiation fields, and therefore it is not
229 an important species in CAM.

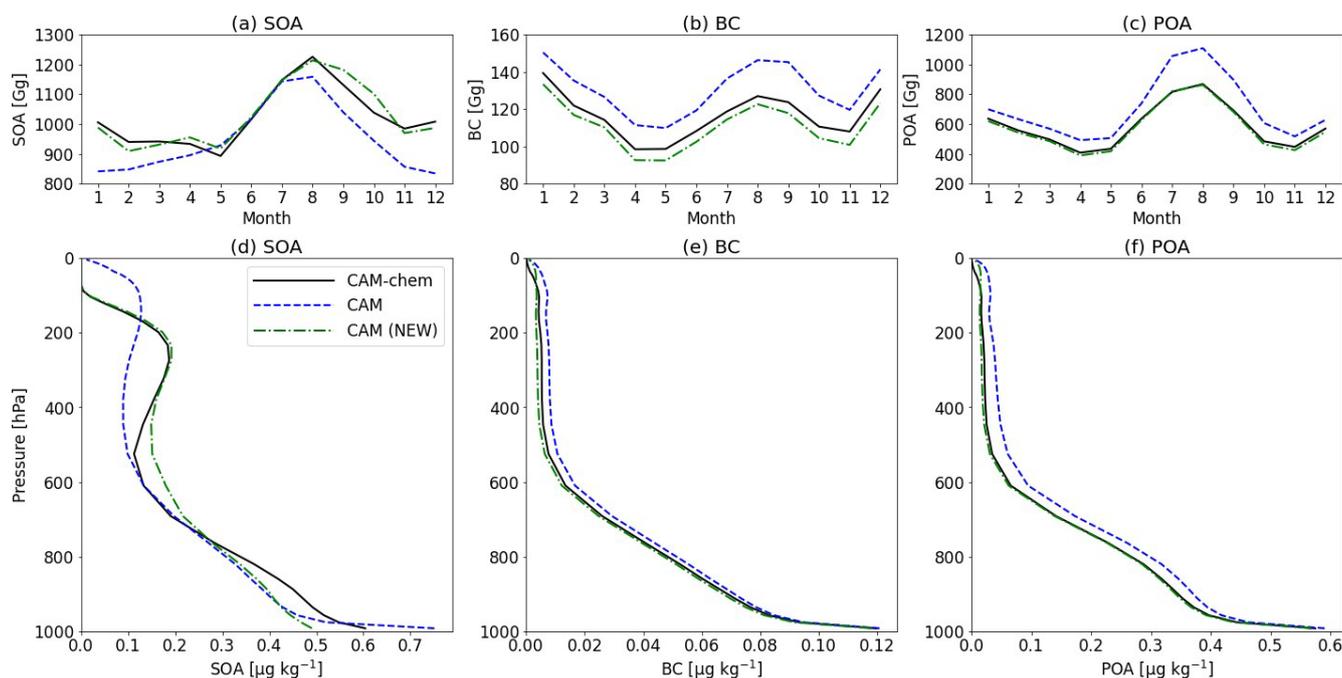


230 **Table 2.** Global annual mean burden of carbonaceous aerosols (SOA, SOAG (SOA in the gas phase),
 231 BC, and POA) and radiation fields (FSNT (net shortwave flux at top of model), FLNT (net longwave
 232 flux at top of model), top of the atmosphere (TOA) imbalance, SWCF (shortwave cloud forcing),
 233 LWCF (longwave cloud forcing)). Because CAM uses the offline biogenic SOA emissions, SOA in the
 234 default CAM is not affected by the CLM option (Sect. 2.5). Units are Gg for aerosols and $W m^{-2}$ for
 235 radiation fields.

Simulation	SOA scheme	SOA	SOAG	BC	POA	FSNT	FLNT	TOA imbalance	SWCF	LWCF
2013 (Nudged)	CAM-chem	1022	484	117	587	236.7	238.7	-2.0	-50.5	22.2
	CAM	948	118	131	704	237.7	239.2	-1.5	-49.6	21.7
	CAM (NEW)	1027	129	111	574	237.3	239.3	-2.0	-49.8	21.6
1850s (SP)	CAM-chem	780	367	31	299	232.3	235.0	-2.7	-54.6	26.3
	CAM	699	102	43	435	233.1	235.4	-2.3	-53.7	25.7
	CAM (NEW)	747	94	30	300	232.7	235.3	-2.7	-54.0	25.7
2000s (SP)	CAM-chem	793	375	89	510	231.3	234.0	-2.7	-56.1	25.7
	CAM	796	102	102	635	232.0	234.4	-2.4	-55.3	25.1
	CAM (NEW)	744	105	83	488	231.7	234.4	-2.7	-55.6	25.1
1850s (BGC)	CAM-chem	826	357	31	302	232.2	235.0	-2.8	-55.0	26.4
	CAM (NEW)	770	89	31	304	232.6	235.3	-2.7	-54.3	25.9
2000s (BGC)	CAM-chem	982	411	88	510	231.3	234.0	-2.7	-56.3	25.8
	CAM (NEW)	952	109	83	490	231.6	234.3	-2.7	-55.8	25.2
2000s - 1850s (SP)	CAM-chem	13	8	57	210	-0.98	-0.97	-0.01	-1.47	-0.54
	CAM	97	0	60	200	-1.15	-1.04	-0.11	-1.67	-0.66
	CAM (NEW)	-3	11	52	188	-0.98	-0.91	-0.07	-1.58	-0.70
2000s - 1850s (BGC)	CAM-chem	156	54	57	208	-0.92	-1.08	0.16	-1.31	-0.59
	CAM (NEW)	182	19	52	185	-0.96	-0.97	0.01	-1.44	-0.75



260 Although two CAM cases show similar global SOA burdens to CAM-chem, their temporal and
261 spatial distributions are very different. Figure 2 shows the monthly timeseries and mean vertical profile
262 of the global SOA burden simulated by CAM and CAM-chem in 2013. The SOA underestimation
263 during the northern hemisphere winter time and the SOA build-up in the upper atmosphere (< 100 hPa)
264 are greatly improved in the new CAM SOA scheme. There is still a discrepancy between the new CAM
265 and CAM-chem such as SOA at around 500 hPa and at the surface (Fig. 2d), due to the limitation of
266 using only one volatility bin in CAM (to reduce the computational cost). One fixed volatility bin with
267 one enthalpy value cannot fully reproduce gas-phase semivolatiles simulated by five bins and the
268 temperature dependency of volatility changes.



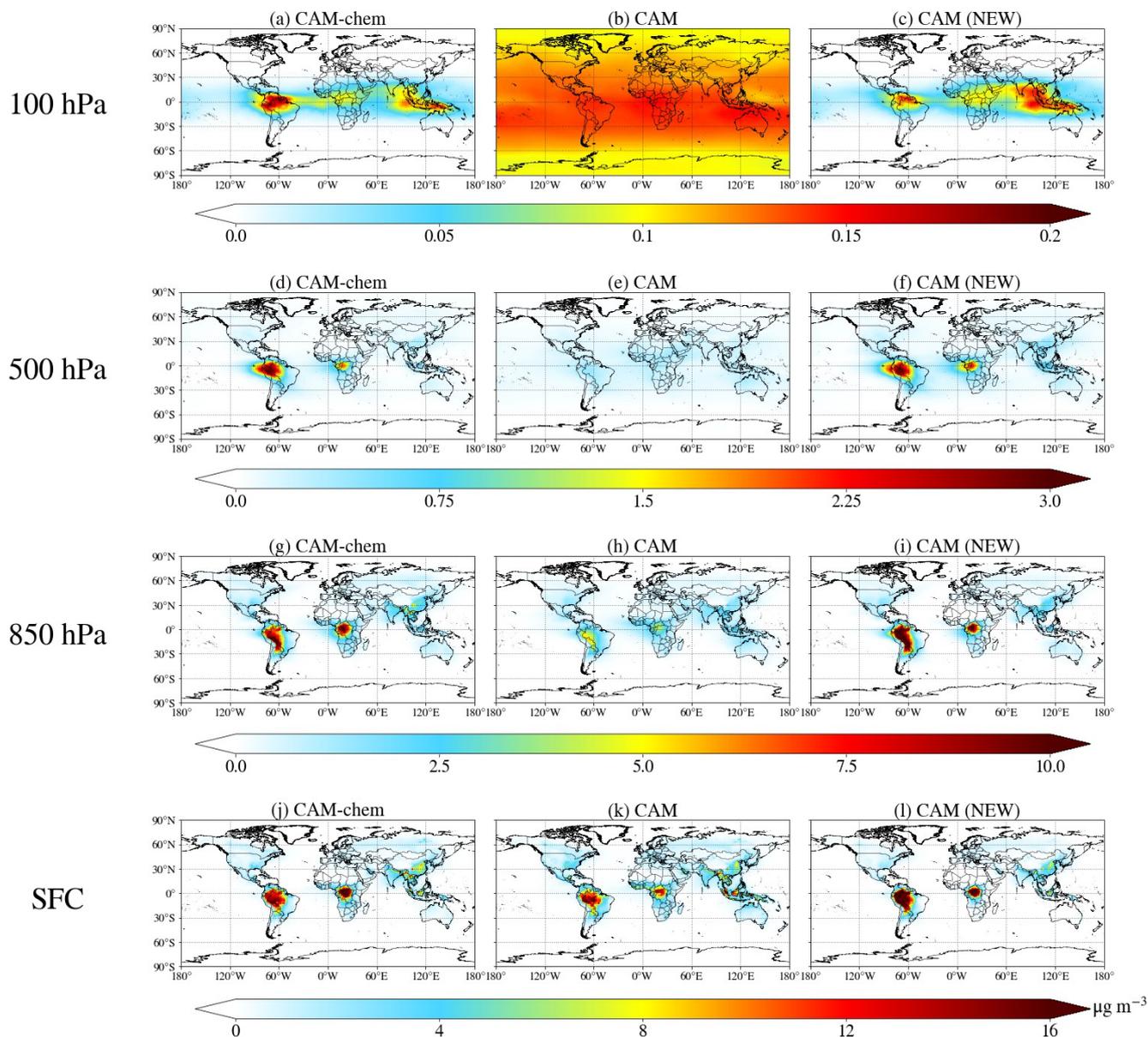
269 **Figure 2.** Monthly timeseries of global atmospheric burden (first row) and vertical distributions (second
270 row) of annual average SOA, BC, and POA simulated by CESM2.



271 Figure 3 shows the global spatial distribution of SOA at 100 hPa, 500 hPa, 850 hPa, and the surface
272 levels simulated by CAM-chem and CAM. In the current CAM simulation, the main source regions
273 (South America and Africa) are well represented at the surface layer (Fig. 3k) but do not appear in the
274 free troposphere and above (panels b, e, and h). This is because the current CAM SOA scheme
275 generates semivolatiles directly from the surface emissions while the CAM-chem SOA scheme needs
276 more time for VOC reactions to make semivolatiles, which can form SOA in the free troposphere. The
277 intermediate tracer (SOAE) in the new CAM implicitly considers this process and successfully captures
278 SOA peaks in the free troposphere (panels c, f, and i).

279 In addition, the current CAM fails to reproduce the sharp gradient of CAM-chem SOA above 200
280 hPa (Fig. 2d) and simulates too much SOA globally (Fig. 3b). The missing loss processes (deposition of
281 semivolatiles and photolytic loss of SOA) and higher temperature dependency (enthalpy) of saturation
282 vapor pressure result in more SOA in the current CAM simulation. This problem is solved in the new
283 CAM SOA scheme (Fig. 3c).

284 Significant improvements are also found for BC and POA. The current CAM simulates up to ~45%
285 differences while the new CAM shows up to ~7% differences for BC and POA (Table 2). Unlike SOA,
286 seasonalities of BC and POA are well represented in the current CAM (panels b and c in Fig. 2), since
287 BC and POA schemes are the same between CAM and CAM-chem. Spatial distributions are also
288 similar (Figs. S2–S5) except for the Arctic regions in the upper atmosphere. This difference can
289 significantly affect the radiation budget in the Arctic region (Sect. 3.2), which should be important for
290 climate studies focusing on the Arctic.

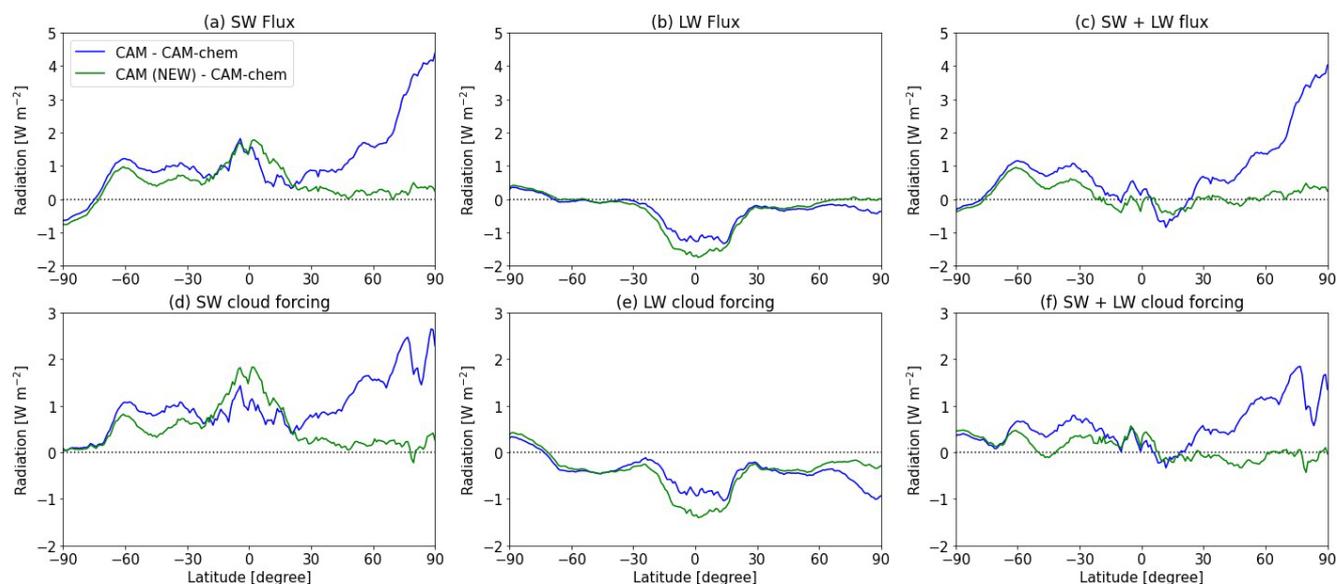


291 **Figure 3.** Global maps of SOA concentrations in 2013 simulated by CAM-chem (first column), CAM
292 (second column), and CAM (NEW) (third column) at four different vertical levels (surface, 850 hPa,
293 500 hPa, and 100 hPa). The difference maps between CAM and CAM-chem are available in Fig. S1.



294 3.2 Radiation fields

295 As aerosols can affect radiative fluxes through direct and indirect effects, here we investigate the
296 radiation changes with the new SOA scheme in CAM, in terms of the difference between CAM and
297 CAM-chem. Figure 4 shows the zonal averages of net shortwave (SW) and longwave (LW) fluxes and
298 cloud forcings in CAM compared to CAM-chem. The most notable differences occur in the high
299 latitudes in the Northern Hemisphere, similar to aerosol concentration changes shown in Sect 3.1. Both
300 aerosol-radiation and aerosol-cloud interactions almost equally contribute to the positive bias (panels a
301 and d). This strong positive bias of the SW flux in the current CAM is greatly improved with the new
302 SOA scheme.



303 **Figure 4.** Zonal averages of the radiation difference in 2013 between CAM and CAM-chem. Radiative
304 fluxes at the top of the model are presented in the first row (a-c) and cloud forcings are shown in the
305 second row (d-f).



306 Biases are slightly increased over the Tropics when it comes to SW and LW fluxes individually,
307 which are mainly caused by the cloud effects as shown in Figs. 4d and 4e, but those are canceled out in
308 terms of the total radiation (Figs. 4c and 4f). Overall, the new SOA scheme shows slight improvements
309 in other latitudes as well in addition to the Arctic region. The low bias can be further confirmed by
310 global spatial distributions shown in Fig. S6, the new CAM simulation shows fewer biases in most of
311 the locations globally (panels h and i in Fig. S6).

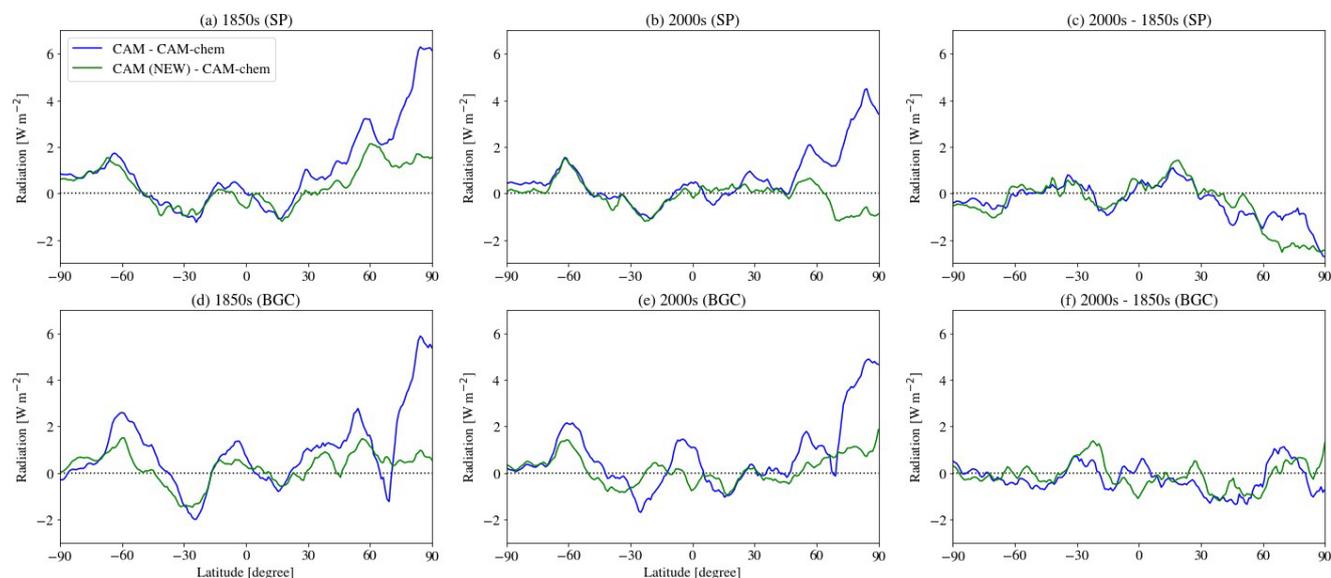
312 **3.3 Historical simulations**

313 Analogous to the simulation results with nudged meteorology in Sect 3.1 and 3.2, the new SOA
314 scheme produces more consistent results with CAM-chem than the current CAM SOA scheme (Table
315 2), especially for BC and POA burdens that are affected by SOA through microphysics. The new SOA
316 scheme also captures the increased SOA burden in the 2000s compared to the 1850s when using the
317 BGC option, which is mainly caused by increased biogenic VOC emissions (Fig. S7).

318 Figure S7 further shows that interannual variability may not be a significant factor for isoprene
319 emissions on a 10-years time scale, but this would be important for climate studies with more than 100
320 years of simulation time (1850s vs 2000s). The offline emissions used in the default CAM have no
321 interannual variability, thus not accounting for emission response to climate change.

322 The high bias of the SW + LW flux over the Arctic from the nudged meteorology simulations (Fig.
323 4c) is also found in all historical simulations as shown in Fig. 5. The default CAM shows the high bias
324 for both 1850s and 2000s simulations. However, in terms of the difference between the 2000s and
325 1850s, the biases cancel out, and as a result, the difference between CAM and CAM-chem becomes
326 small (Figs. 5c and 5f). This cancellation implies that previous CAM studies focusing on radiative
327 forcing are still valid, as radiative forcing is calculated as present minus preindustrial radiative effects.

328 In terms of global averages (Table 2), the new CAM also shows improvements, especially for
329 shortwave radiation, not only for absolute values but also for the difference between present and
330 pre-industrial simulations.



331 **Figure 5.** Zonal averages of the SW + LW flux difference in historical simulations (1850s (a and d),
332 2000s (b and e), and 2000s - 1850s (c and f)) between CAM and CAM-chem. Note that the results from
333 the default CAM simulations are the same for SP and BGC because the default CAM uses offline
334 biogenic emissions. Only CAM-chem results affect the difference between SP and BGC simulations
335 (blue lines).

336 4 Conclusion and possible future developments of the aerosol scheme in CAM

337 In this study, we developed a new SOA scheme for use in CAM with simple chemistry. This new
338 SOA scheme was designed to close the gap between CAM and CAM-chem in terms of aerosols and
339 radiative effects while maintaining computational efficiency. The new SOA scheme was derived based
340 on the parameters used in the VBS scheme in CAM-chem, without changing the overall architecture of
341 the simple SOA scheme in the current CAM. As a result, the computational cost remained almost the
342 same with the new SOA scheme (within the range of computing environment variability).

343 CAM simulation results with the current and new SOA schemes were investigated in terms of
344 carbonaceous aerosols and radiative fluxes. There was no significant bias in terms of the global SOA
345 burden of the current CAM SOA scheme because it was tuned by increasing SOA emissions by 50%
346 (Liu et al., 2012). However, the current CAM SOA scheme was insufficient to reproduce the temporal



347 and spatial (both horizontally and vertically) variabilities of CAM-chem SOA, and the new SOA
348 scheme improved the performance of these variabilities.

349 The new SOA scheme also improved the simulation of other carbonaceous aerosols (BC and POA)
350 through the microphysics of MAM4. Since BC and POA emissions are the same for all model cases and
351 those aerosols are chemically inert, temporal and horizontal spatial variabilities are generally similar to
352 each other but the absolute concentrations were improved when using the new SOA scheme. The
353 overestimation of BC in CAM was greatly reduced compared to CAM-chem, from ~45% in the current
354 SOA scheme to ~7% in the new SOA scheme. POA is also improved in the same manner. Major
355 improvements were made in the Arctic region for aerosol concentrations in the free troposphere and
356 above.

357 The improvements in simulating aerosol fields led to more consistent radiative fluxes between
358 CAM and CAM-chem, especially over the high-latitude regions in the Northern Hemisphere. The SW +
359 LW flux at the top of the model was different by up to 6 W m^{-2} and it is persistent regardless of the
360 simulation periods. However, in terms of the radiative forcing which is calculated from the difference
361 between present and pre-industrial conditions, both the current and new CAM simulations showed no
362 significant differences. While the studies investigating the instantaneous radiative effects will need to
363 use the new SOA scheme, the current SOA scheme would still be valid for studies focusing on radiative
364 forcing.

365 On the practical side, the new SOA scheme has advantages in keeping up with the updates, as it
366 uses the same precursor emissions as the VBS scheme in CAM-chem. The new SOA scheme uses
367 online biogenic emissions as CAM-chem does, therefore the difference between SP and BGC options
368 can be calculated for SOA. If there is a future update in the VBS scheme in CAM-chem, the
369 corresponding updates in CAM can be done easily by changing the namelist file.

370 Although significant advances have been made in SOA concentration simulation in this study, the
371 aerosol module in CAM still has room for further development. Currently, CAM reads the offline
372 monthly oxidant fields simulated by CAM-chem but oxidants such as OH and O₃ have strong diurnal
373 variations. It would not be computationally feasible for CAM to calculate or read oxidants every hour,
374 but applying constant diurnal profile values to the monthly fields would not add significant



375 computational costs. It may be important for SO₂ oxidation and sulfate formation as well. The formation
376 of SOAG from SOAE is calculated using a 1-day lifetime, but future versions could use the reaction
377 rate constant with OH if the diurnal variation of oxidant fields is introduced in CAM.

378 The new SOA scheme can be further adjusted, for example, for studies focusing on surface aerosol
379 fields, users can easily change SOA yields for different emission sources through namelist changes.
380 Vertical shapes can be also adjusted by changing the parameters such as the enthalpy of vaporization,
381 saturation vapor pressure, and photolysis rates in the future.

382 **Code and data availability.** CESM is an open-source community model and is publicly available at:
383 <https://github.com/ESCOMP/CESM>. The new SOA scheme is included in the development version of
384 the CAM (<https://github.com/ESCOMP/CAM>, tag name: cam6_3_093) and will be publicly available in
385 the next CESM release. The model results used in this study are available on the NCAR Digital Asset
386 Service Hub (DASH) at TBD.

387 **Author contributions.** DSJ, ST, and LKE designed the research and developed the SOA scheme. SW
388 developed the OASISS scheme. DSJ, ST, and FV conducted CESM simulations. DSJ wrote the
389 manuscript. All authors contributed to editing the manuscript.

390 **Competing interests.** The authors declare that they have no conflict of interest.

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