



1 **Adjoint of GEOS-Chem model by supporting HEMCO emission**
2 **inventories and MERRA-2 meteorological data**

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11

12 **Abstract**

13 Adjoint of the GEOS-Chem model has been widely used to constrain the sources of various
14 atmospheric pollutants. Here we provide an updated version (GC-Adjoint-HEMCO) of the
15 adjoint of GEOS-Chem model to support the MERRA-2 meteorological data and Harmonized
16 Emissions Component (HEMCO) emission inventories. State-of-the-art inventories, such as
17 CEDS (Community Emissions Data System), MIX, NEI2011 (National Emissions Inventory),
18 and GFED4 (Global Fire Emission Database), are supported in GC-Adjoint-HEMCO. We find
19 good agreements in carbon monoxide (CO) emissions from various inventories, chemical
20 sources and sinks, and surface and column CO concentrations between GC-Adjoint-HEMCO
21 and GEOS-Chem (v12-8-1) forward simulations. Furthermore, observing system simulation
22 experiments (OSSE) are employed to evaluate the performance of GC-Adjoint-HEMCO in 4D
23 variational (4D-var) assimilations. We find underestimations by approximately 15% in the a
24 posteriori anthropogenic CO emissions over North America and Europe due to limited
25 coverage of observations by smoothing the pseudo-CO observations with Measurement of
26 Pollution in the Troposphere (MOPITT) averaging kernels. As an example application of GC-
27 Adjoint-HEMCO, we constrain anthropogenic CO emissions in 2015 by assimilating MOPITT
28 CO observations. The a posteriori anthropogenic CO emission estimates derived in this work
29 match well with Jiang et al. (2017) in North America and Africa but are overestimated in Asia,
30 South America and Australia, which could be associated with the different treatment of



31 MOPITT CO observations over ocean grids and the large differences in CO chemical sources
32 and sinks. The updated model developed in this work is a useful extension of the adjoint of
33 GEOS-Chem model.

34

35 **1. Introduction**

36 GEOS-Chem is a global 3D chemical transport model (CTM) and has been widely used
37 to analyze the sources and variabilities of various atmospheric compositions (Li et al., 2019;
38 Wang et al., 2019; Hammer et al., 2020; Jiang et al., 2022). GEOS-Chem model is driven by
39 meteorological reanalysis data from the Goddard Earth Observing System (GEOS) of the
40 Global Modeling and Assimilation Office (GMAO). Emissions in GEOS-Chem model are
41 calculated based on the Harmonized Emissions Component (HEMCO) (Keller et al., 2014; Lin
42 et al., 2021) with default support for state-of-the-art inventories such as CEDS (Community
43 Emissions Data System) (Hoesly et al., 2018), MIX (Li et al., 2017) and NEI2011 (National
44 Emissions Inventory). Based on GEOS-Chem forward model, the adjoint of the GEOS-Chem
45 model (Henze et al., 2007) further provides the capability of backward simulation of physical
46 and chemical processes within the 4D variational (4D-var) framework. The major advantage
47 of the adjoint model is obtaining the sensitivity of atmospheric concentrations to multiple
48 model variables within a single backward simulation. The major applications of the adjoint of
49 GEOS-Chem model include inverse analyses of atmospheric pollutant emissions by
50 minimizing the difference between simulations and observations (Jiang et al., 2017; Zhang et
51 al., 2018; Qu et al., 2022) as well as sensitivity analyses to analyze the sources of atmospheric
52 pollutants (Jiang et al., 2015; Zhao et al., 2019; Dedoussi et al., 2020).

53 The algorithm of the 4D-var framework requires identical model processes in the forward
54 and backward simulations. Ideally, the code for the backward simulation should be updated
55 following the GEOS-Chem forward code to take advantage of the new features in GEOS-Chem



56 forward simulations. However, the updates in the backward code are difficult and usually
57 delayed. For example, the MERRA-2 meteorological reanalysis data with temporal coverage
58 of 1979-present were supported in the GEOS-Chem forward simulations in v11-01. The adjoint
59 of GEOS-Chem model does not support MERRA-2, and thus, long-term analysis must combine
60 different meteorological reanalysis data, such as GEOS-4 (1985-2007), GEOS-5 (2004-2012)
61 and GEOS-FP (2012-present), which can lead to discontinuity in the derived trends of
62 emissions (Jiang et al., 2017). Furthermore, the HEMCO emission module was included in the
63 GEOS-Chem forward simulations in v10-01. Adjoint of GEOS-Chem model does not support
64 the HEMCO module, and thus, updated emission inventories such as CEDS, MIX and NEI2011
65 cannot be read conveniently, which can affect the performance in the simulated atmospheric
66 compositions.

67 In this work, we develop the capability of the adjoint of GEOS-Chem model to support
68 MERRA-2 meteorological data and HEMCO emission inventories. The results presented in
69 this paper show the development, evaluation, and application of the developed capability to
70 constrain carbon monoxide (CO) emissions by assimilating CO measurements from the
71 Measurement of Pollution in the Troposphere (MOPITT). CO is one of the most important
72 atmospheric pollutants and plays a key role in tropospheric chemistry. Sources of atmospheric
73 CO include fossil fuel combustion, biomass burning and oxidation of hydrocarbons. The major
74 sink of atmospheric CO is hydroxyl radical (OH). The simple chemical sink of atmospheric
75 CO allows us to simulate atmospheric CO with linearized chemistry; for example, the tagged-
76 CO mode of the GEOS-Chem model can reduce the calculation cost by 98% with respect to
77 the full chemistry mode by reading archived monthly OH fields. The tagged-CO mode of the
78 GEOS-Chem model has been widely used to investigate the sources and variabilities of
79 atmospheric CO in recent decades (Heald et al., 2004; Kopacz et al., 2009; Jiang et al., 2017).
80 The capability presented in this work is thus based on the tagged-CO mode, as it can effectively



81 accelerate the model development.

82 This paper is organized as follows: in Section 2, we describe the adjoint of GEOS-Chem
83 model, the development of the new capability to support the MERRA-2 meteorological data
84 and HEMCO emission inventories, and the MOPITT CO observations used in this work. In
85 Section 3, we evaluated the performance of the developed model and presented an example
86 application of 4D-var assimilation to constrain anthropogenic CO emissions in 2015 by
87 assimilating MOPITT CO observations. Our conclusions follow in Section 4.

88

89 **2. Methodology and Data**

90 **2.1 Adjoint of the GEOS-Chem model**

91 We use version v35n of the adjoint of GEOS-Chem model. Our analysis is conducted at
92 a horizontal resolution of $4^{\circ} \times 5^{\circ}$ with 47 vertical levels and employs the CO-only simulation
93 (tagged-CO mode). The global default anthropogenic emission inventory in the standard
94 version of the adjoint of GEOS-Chem model (hereafter referred to as GC-Adjoint-STD) is
95 Global Emissions Initiative (GEIA), but is replaced by the following regional emission
96 inventories: NEI2008 in North America, the Criteria Air Contaminants (CAC) inventory for
97 Canada, the Big Bend Regional Aerosol and Visibility Observational (BRAVO) Study
98 Emissions Inventory for Mexico (Kuhns et al., 2003), the Cooperative Program for Monitoring
99 and Evaluation of the Long-range Transmission of Air Pollutants in Europe (EMEP) inventory
100 for Europe in 2000 (Vestreng and Klein, 2002) and the INTEX-B Asia emissions inventory for
101 2006 (Zhang et al., 2009). Biomass burning emissions are based on the GFED3 (van der Werf
102 et al., 2010).

103 The objective of the 4D-var approach is to minimize the difference between simulations
104 and observations described by the cost function by adjusting the CO emissions iteratively
105 (Henze et al., 2007):



$$J(x) = \sum_{i=1}^N [F_i(x) - z_i]^T S_{\Sigma}^{-1} [F_i(x) - z_i] + (x - x_a)^T S_a^{-1} (x - x_a) \quad (1)$$

106 where x is the state vector of CO emissions, N is the number of observations that are
107 distributed in time over the assimilation period, z_i is a given measurement, and $F(x)$ is the
108 forward model. The error estimates are assumed to be Gaussian and are given by S_{Σ} , the
109 observational error covariance matrix, and S_a , the a priori error covariance matrix. We assume
110 a uniform observation error of 20%. The combustion CO sources (fossil fuel, biofuel and
111 biomass burning) and the oxidation source from biogenic volatile organic compound (VOC)
112 are combined, assuming a 50% uniform a priori error. We optimize the source of CO from the
113 oxidation of methane (CH_4) separately as an aggregated global source, assuming an a priori
114 uncertainty of 25%. The CO emission estimates are optimized with monthly temporal
115 resolution. We performed 40 iterations (forward + backward simulations) for each month, and
116 the a posteriori CO emission estimates were calculated based on the last accepted iteration,
117 which usually corresponded to the iteration with the lowest cost function.

119 2.2 Updates in emission inventories

120 A major objective of this work is to obtain consistent emissions between the adjoint of
121 GEOS-Chem model and the GEOS-Chem forward model (v12-08-01, hereafter referred to as
122 GC-v12). As shown in Fig. 1, we first initialize the array in [INITIAL] and batch read the
123 emission data in [READ_DATA], which were interpolated offline with $1^\circ \times 1^\circ$ resolution by
124 considering the mass conservation. Here, the data include the emission inventory data listed in
125 Table S1 (see the SI), the corresponding scaling factor data and the mask map files of domain
126 definitions. The data are scaled in [SCALE_DATA] by multiplying the corresponding annual,
127 season, month, week, and 24-hour emission factors and are then online interpolated to the
128 current resolution ($4^\circ \times 5^\circ$ in this work) of the model by [RGRID_DATA], which was followed
129 by the application of region masks in [MASK].

130 The emission variable of CO obtained in this part is written to the model memory in



131 emission.f and emission_adj.f. by calling DO_EMISSIONS. The GET_[TRACER] subroutines
132 are used to obtain the CO emission variable, which participates in the calculation of
133 physicochemical processes in the model, to interact with other modules. Finally, the variable
134 is cleaned from the memory by the [CLEANUP] module. It should be noted that a two-step
135 interpolation is employed in GC-Adjoint-HEMCO following GC-Adjoint-STD, for example,
136 $0.1^\circ \times 0.1^\circ$ to $1^\circ \times 1^\circ$ and then to $4^\circ \times 5^\circ$ for the NEI2011 inventory, which is different from the
137 one-step interpolation in GC-v12, for example, $0.1^\circ \times 0.1^\circ$ to $4^\circ \times 5^\circ$ directly for the NEI2011
138 inventory. The different interpolation methods can lead to differences in the interpolated
139 emission data.

140 As shown in Table S1, the CEDS emission inventory ($0.5^\circ \times 0.5^\circ$) is adopted in GC-
141 Adjoint-HEMCO to provide global default emissions for 1750-2019. The diurnal scale factors
142 are applied to obtain CO emissions at different moments of the day. Fig. S1 (see the SI) shows
143 CEDS CO emissions in 2015 in GC-v12 and GC-Adjoint-HEMCO and GEIA CO emissions
144 in GC-Adjoint-STD, and we find noticeable differences in CO emissions between CEDS and
145 GEIA. As shown in Table 1, the CEDS CO emissions in 2015 were 613.57 and 613.85 Tg/y in
146 GC-v12 and GC-Adjoint-HEMCO, respectively, with a relative difference of 0.05% between
147 GC-v12 and GC-Adjoint-HEMCO. The GEIA CO emissions in 2015 were 445.88 Tg/year in
148 GC-Adjoint-STD.

149 The default CEDS inventory is replaced by the following regional emission inventories
150 in GC-Adjoint-HEMCO: MIX in Asia ($0.25^\circ \times 0.25^\circ$), NEI2011 in the United States
151 ($0.1^\circ \times 0.1^\circ$), DICE_AFRICA and EDGARV43 in Africa ($0.1^\circ \times 0.1^\circ$) and APEI in Canada
152 ($0.1^\circ \times 0.1^\circ$). As shown in Fig. S2 (see the SI), the MIX inventory provides Asian emissions in
153 2008-2010, accompanied by diurnal scale factors to describe daily emission variation. The
154 $1^\circ \times 1^\circ$ scale factors in the AnnualScalar.geos.1x1.nc file further provide the annual variation in
155 1985-2010. As shown in Table 1, the MIX CO emissions in 2015 were 321.18 and 321.71 Tg/y



156 in GC-v12 and GC-Adjoint-HEMCO, respectively, with a relative difference of 0.17% between
157 GC-v12 and GC-Adjoint-HEMCO. The INTEX-B CO emissions in 2015 were 353.03 Tg/y in
158 GC-Adjoint-STD.

159 The NEI2011 inventory (Fig. S3, see the SI) provides anthropogenic emissions for the
160 United States in 2011 with annual scalar factors from 2006-2013. The mid-week and weekend
161 factors are read from NEI99.dow.geos.1x1.nc file since 1999 with all CO factors of 1.0 on
162 weekdays and between 0.990 and 0.997 on Saturdays and Sundays. The NEI2011 CO
163 emissions in 2015 were 35.83 and 37.70 Tg/y in GC-v12 and GC-Adjoint-HEMCO,
164 respectively, with a relative difference of 5.22% between GC-v12 and GC-Adjoint-HEMCO.
165 The NEI2008 CO emissions in 2015 were 52.87 Tg/y in GC-Adjoint-STD. APEI (Fig. S4, see
166 the SI) is the primary source of anthropogenic emissions in the Canadian domain. The APEI
167 CO emissions in 2015 were 6.10 and 6.17 Tg/y in GC-v12 and GC-Adjoint-HEMCO,
168 respectively, with a relative difference of 1.14% between GC-v12 and GC-Adjoint-HEMCO.
169 The CAC CO emissions in 2015 were 10.20 Tg/y in GC-Adjoint-STD. Following GC-v12, the
170 CO emissions in APEI are enhanced by 19% to account for coemitted VOC in the tagged-CO
171 simulation.

172 Emissions for the African domain are provided by the combination of DICE_AFRICA
173 and EDGARV43 (Fig. S5, see the SI). Here DICE_AFRICA includes anthropogenic and
174 biofuel emissions in 2013. We read the DICE_AFRICA emissions data into the model in two
175 types according to the guidelines of the inventory. Emissions from sectors such as automobiles
176 and motorcycles are aggregated into anthropogenic sources, and household-generated
177 emissions such as charcoal and agricultural waste are aggregated into biofuel sources. Efficient
178 combustion emissions from EDGAR v4.3 in 1970-2010 then compensate for the lacking
179 sources in DICE_AFRICA. Daily variation factors for CO are also used here for emissions
180 across the African region. The 2010 CO seasonal scale factors are used in EDGAR v4.3 for



181 sectoral emission sources. The DICE_AFRICA and EDGARV43 CO emissions in 2015 were
182 83.42 and 83.02 Tg/y in GC-v12 and GC-Adjoint-HEMCO, respectively, with a relative
183 difference of -0.48% between GC-v12 and GC-Adjoint-HEMCO. Following GC-v12, the CO
184 emissions in DICE_AFRICA and EDGARV43 are enhanced by 19% to account for coemitted
185 VOC in the tagged-CO simulation.

186 The biomass burning emission inventory in the GC-Adjoint-HEMCO is GFED4 (Fig.
187 S6, see the SI), which includes dry matter emissions from a total of seven sectors in 1997-2019.
188 The same GFED_emission_factors.H header file as in the GC-v12 version is read in the GC-
189 Adjoint-HEMCO. This file contains the ratio factors of atmospheric pollutants, and we
190 multiply the ratio factors one by one according to the ID of each species to ensure that the
191 species in the model have biomass burning sources. The GFED4 CO emissions in 2015 were
192 437.13 and 435.89 Tg/y in GC-v12 and GC-Adjoint-HEMCO, respectively, with a relative
193 difference of -0.28% between GC-v12 and GC-Adjoint-HEMCO. The GFED3 CO emissions
194 in 2015 were 382.04 Tg/year in GC-Adjoint-STD. Following GC-v12, the combustion CO
195 sources in biomass burning are enhanced by 5% to consider the CO generated by VOC in the
196 tagged-CO simulation.

197 Fig. 2 shows the total combustion CO emissions in 2015 from GC-v12, GC-Adjoint-
198 HEMCO and GC-Adjoint-STD. As shown in Table 2, the regional combustion CO emissions
199 are 320.66 and 320.38 Tg/y (Asia), 73.96 and 66.93 Tg/y (North America), 199.51 and
200 193.29/y Tg (Africa), 79.04 and 78.91 Tg/y (South America), 31.58 and 30.96 Tg/y (Europe)
201 and 12.24 and 11.99 Tg/y (Australia) in GC-v12 and GC-Adjoint-HEMCO, respectively. Fig.
202 3 further shows the monthly combustion CO emissions in 2015 from GC-v12, GC-Adjoint-
203 HEMCO and GC-Adjoint-STD, and there are good agreements in the monthly variation of CO
204 emissions between GC-v12 and GC-Adjoint-HEMCO. The CO emissions in GC-Adjoint-STD
205 are similar to those in GC-v12 and GC-Adjoint-HEMCO in winter and spring but with large



206 differences in summer and autumn. This seasonal difference may reflect the influence of
207 different emission inventories on biomass burning.

208 **2.3 Updates in CO chemical sources and sinks**

209 The biogenic emissions in GC-Adjoint-STD are Model of Emissions of Gases and
210 Aerosols from Nature, version 2.0 (MEGANv2.0, Guenther et al. (2006)) in the full chemistry
211 simulation but are GEIA in the tagged-CO simulation (Fig. S7, see the SI). Fisher et al. (2017)
212 demonstrated improvement in modeled CO concentrations in tagged-CO simulation by reading
213 archived VOC- and CH₄-generated CO fields provided by full chemistry simulation. The
214 archived VOC- and CH₄-generated CO fields in 2013 (PCO_3Dglobal.geosfp.4x5.nc) were set
215 as the default CO chemical sources in the tagged-CO simulation in GC-v12 and supported in
216 GC-Adjoint-HEMCO. As shown in Table 2, the CO chemical sources (columns) obtained by
217 reading the archived VOC- and CH₄-generated CO fields demonstrate good agreement between
218 GC-v12 and GC-Adjoint-HEMCO. However, they are 30-60% lower than those in GEIA in
219 GC-Adjoint-STD, and this difference could be partially associated with the inconsistency
220 between the archived VOC-generated CO fields in 2013 and the actual meteorological data in
221 2015 in the simulation.

222 The default CH₄-generated CO emissions in GC-Adjoint-STD (Fig. S8, see the SI) are
223 calculated based on averaged CH₄ concentrations in four latitude bands (90°S - 30°S, 30°S -
224 00°S, 00°N - 30°N, 30°N - 90°N), which are based on Climate Monitoring and Diagnostics
225 Laboratory (CMDL) surface observations and Intergovernmental Panel on Climate Change
226 (IPCC) future scenarios. As shown in Table 2, there are good agreements in the CH₄-generated
227 CO emissions between GC-v12 and GC-Adjoint-HEMCO by reading
228 PCO_3Dglobal.geosfp.4x5.nc, and they are 20-60% lower than those in CMDL/IPCC in GC-
229 Adjoint-STD. Furthermore, the default archived monthly OH fields were updated following
230 GC-v12 with updated calculations for the decay rate (KRATE, from JPL 03 to JPL 2006) in



231 GC-Adjoint-HEMCO. The subsequent CO sinks (Fig. S9, see the SI) in GC-v12 and GC-
232 Adjoint-HEMCO are 20-40% higher than those in GC-Adjoint-STD.

233 **2.4 Updates in meteorological data**

234 The MERRA-2 meteorological data (1979-present) are supported in GC-Adjoint-
235 HEMCO to ensure long-term consistency in the meteorological data in inverse analyses. The
236 code porting to support MERRA-2 is more direct than emission inventories because the
237 meteorological variables and vertical resolutions of MERRA-2 are the same as those of GEOS-
238 FP (2012-present), while GEOS-FP is already supported by GC-Adjoint-STD. Fig. 4A-B show
239 the averages of surface CO concentrations in 2015 from GC-Adjoint-HEMCO driven by
240 MERRA-2 and GEOS-FP, respectively. Our results demonstrate lower surface CO
241 concentrations driven by MERRA-2 (Fig. 4C), although there is good agreement in the spatial
242 distributions of CO concentrations. Similarly, Fig. 4D-F show the averages of CO columns in
243 2015 from GC-Adjoint-HEMCO driven by MERRA-2 and GEOS-FP and their differences.
244 Despite the noticeable differences in surface CO concentrations (Fig. 4C), the differences in
245 CO columns (Fig. 4F) are much smaller, and the modeled CO columns driven by MERRA-2
246 are higher than those driven by GEOS-FP over the Indian Ocean. The discrepancy between
247 surface and column CO in Fig. 4 may reflect the impacts of different convective transports on
248 the modeled CO concentrations.

249 **2.5 MOPITT CO measurements**

250 The MOPITT data used here were obtained from the joint retrieval (V7J) of CO from
251 thermal infrared (TIR, 4.7 μ m) and near-infrared (NIR, 2.3 μ m) radiances using an optimal
252 estimation approach (Worden et al., 2010; Deeter et al., 2017). The retrieved volume mixing
253 ratios (VMR) are reported as layer averages of 10 pressure levels with a footprint of 22 km \times
254 22 km. Following Jiang et al. (2017), we reject MOPITT data with CO column amounts less
255 than 5×10^{17} molec/cm² and with low cloud observations. Since the NIR channel measures



256 reflected solar radiation, only daytime data are considered.

257

258 **3. Model evaluation and application**

259 **3.1 CO concentrations in the forward simulations**

260 We first evaluate the performance of GC-Adjoint-HEMCO in forward simulations. Fig.
261 5 shows the averages of surface and column CO concentrations in 2015 from GC-v12, GC-
262 Adjoint-HEMCO and GC-Adjoint-STD. As shown in Table 2, the regional differences between
263 GC-v12 and GC-Adjoint-HEMCO are 2.6%, -5.7%, -4.6%, -1.7%, -1.4% and -3.6% in surface
264 CO concentrations, and -2.3%, -3.6%, -3.3%, -3.1%, -3.3% and -4.1% in CO columns over
265 Asia, North America, Africa, South America, Europe, and Australia, respectively. The
266 differences in CO concentrations between GC-v12 and GC-Adjoint-HEMCO are expected in
267 view of the comparable differences in regional emissions, chemical sources and sinks, as shown
268 in Table 2. The regional differences in CO concentrations between GC-v12 and GC-Adjoint-
269 STD are larger: 4.6%, -10.1%, 6.3%, 22.5%, 6.4% and 25.7% in surface CO concentrations,
270 and -0.7%, -9.9%, 2.5%, 8.0%, -5.8% and 8.5% in CO columns over Asia, North America,
271 Africa, South America, Europe, and Australia, respectively.

272 **3.2 Observing system simulation experiments with pseudo-CO observations**

273 Here we further evaluate the performance of GC-Adjoint-HEMCO as a 4D-var
274 framework using observing system simulation experiments (OSSE). OSSE is a useful method
275 and has been widely used to evaluate the performance of various data assimilation systems
276 (Jones et al., 2003; Barré et al., 2015; Shu et al., 2022). In contrast to assimilations by
277 assimilating actual atmospheric observations, pseudo-observations are usually generated by
278 model simulations and then assimilated in OSSE. The true atmospheric states are known in
279 OSSEs as they are used to produce the pseudo-observations, and consequently, the difference
280 between assimilated and true atmospheric states describes the capability of the assimilation



281 systems to converge to the true atmospheric states in assimilations when assimilating actual
282 observations.

283 The pseudo-observations in this work are produced by archiving CO concentrations from
284 GC-Adjoint-HEMCO forward simulations with the CO emissions unchanged (i.e., the default
285 CO emission inventory such as CEDS, MIX and NEI2011). According to the usage of pseudo-
286 observations, two types of OSSE are performed in this work: 1) full modeled CO fields are
287 assimilated as pseudo-observations so that we have pseudo-CO observations at every grid/level
288 and time step (hereafter referred to as OSSE-FulOBS). This experiment is designed to evaluate
289 the performance of the assimilation system under ideal conditions with full coverage of
290 observations. 2) The modeled CO fields are sampled at the locations/times of MOPITT CO
291 observations and smoothed with MOPITT a priori concentrations and averaging kernels to
292 produce MOPITT-like pseudo-CO observations (hereafter referred to as OSSE-MOPITT). This
293 experiment is designed to evaluate the performance of the assimilation system under actual
294 conditions with limited coverage of observations.

295 In the inverse analysis with the pseudo-CO observations, we reduce the anthropogenic
296 CO emissions by 50% so that the objective of the OSSE is to produce scaling factors that can
297 return the source estimate to the default emissions (i.e., scaling factors of 1.0). Fig. 6A shows
298 the annual scaling factors in 2015 in OSSE-FulOBS. After 40 iterations, the a posteriori
299 anthropogenic CO emission estimates converge to the true states in all major emission regions.
300 As shown in Table 3, the regional scaling factors of OSSE-FulOBS are 1.00, 0.97, 0.97, 1.00,
301 0.98 and 0.94 for anthropogenic CO emissions over Asia, North America, Africa, South
302 America, Europe, and Australia, respectively.

303 Furthermore, Fig. 6D shows the annual scaling factors in OSSE-MOPITT, which are
304 noticeably worse than those in Fig. 6A. The regional scaling factors of OSSE-MOPITT are
305 1.04, 0.88, 1.01, 1.02, 0.84 and 0.81 for anthropogenic CO emissions over Asia, North



306 America, Africa, South America, Europe, and Australia, respectively. With respect to OSSE-
307 FullOBS, the limited coverage of observations in OSSE-MOPITT has resulted in
308 approximately 15% underestimations in the a posteriori CO emission estimates over North
309 America and Europe. In addition, Fig. 6B-C and Fig. 6E-F show the a priori and a posteriori
310 biases in the modeled CO columns. We find dramatic improvements in the modeled CO
311 columns, which confirms the reliability of the 4D-var assimilation system. The difference
312 between Fig. 6B and 6E reflects the influence of the application of MOPITT averaging kernels,
313 which lead to larger negative biases in the a priori simulation. It should be noted that we cannot
314 expect comparable improvement in the actual assimilations because of the potential effects of
315 model and observation errors.

316 **3.3 Anthropogenic CO emissions constrained with MOPITT CO observations**

317 As an example of the application of GC-Adjoint-HEMCO, here we constrain
318 anthropogenic CO emissions in 2015 by assimilating MOPITT CO observations. Fig. 7A shows
319 the relative differences between modeled and MOPITT CO columns at the beginning of each
320 month in 2015 (i.e., biases in monthly initial CO conditions) in the original GEOS-Chem
321 simulations. We find dramatic underestimations in the modeled CO columns by approximately
322 30-40%. As indicated by previous studies (Jiang et al., 2013; Jiang et al., 2017), the biases in
323 monthly initial CO conditions are caused by model biases in CO concentrations accumulated
324 in previous months. A lack of consideration of the model biases, as shown in Fig. 7A, can result
325 in overestimations in the derived monthly CO emission estimates.

326 Following Jiang et al. (2017), a suboptimal sequential Kalman filter (Todling and Cohn,
327 1994; Tang et al., 2022) was employed in this work to optimize the modeled CO concentrations
328 with an hourly resolution by combining GC-Adjoint-HEMCO forward simulation and
329 MOPITT CO observations. The CO concentrations provided by the Kalman filter assimilations
330 were archived at the beginning of each month, which were used as the optimized monthly initial



331 CO conditions in the inverse analysis. As shown in Fig. 7B, the biases in the modeled CO
332 columns in the optimized initial CO conditions are pronounced lower than those in the original
333 simulation (Fig. 7A). The optimization of the initial CO conditions is essential for our inverse
334 analysis, as it can ensure that the adjustments in CO emissions are dominated by the differences
335 between simulations and observations in the current month instead of the 30-40%
336 underestimations in CO columns accumulated in previous months.

337 Fig. 8A shows the distribution of a priori anthropogenic CO emissions in 2015. The
338 regional a priori anthropogenic CO emissions (as shown in Table 4) are 243.53, 34.42, 23.24,
339 30.39, 25.94 and 2.02 Tg/y over Asia, North America, Africa, South America, Europe, and
340 Australia, respectively. As shown in Fig. 8B, our inverse analysis suggests a wide distribution
341 of underestimations in the a priori anthropogenic CO emissions in 2015 except in E. China.
342 The regional scaling factors (Table 4) are 1.16, 1.47, 1.52, 1.41, 1.60 and 1.38, and the a
343 posteriori anthropogenic CO emissions are 283.20, 50.47, 35.34, 42.92, 41.62 and 2.79 Tg/y
344 over Asia, North America, Africa, South America, Europe, and Australia, respectively. As
345 shown in Fig. 8C, we find noticeable underestimations in the modeled CO columns in the a
346 priori simulations, despite the negative biases being much weaker than those in Fig. 7A due to
347 the optimization of the initial CO conditions. The negative biases are effectively reduced in the
348 a posteriori simulation driven by the a posteriori CO emission estimates (Fig. 8D).

349 Finally, we compare the a posteriori CO emission estimates in this work with Jiang et al.
350 (2017), who constrained CO emissions in 2001-2015 with GC-Adjoint-STD by assimilating
351 the same MOPITT CO observations. As shown in Table 4, the a posteriori anthropogenic CO
352 emission estimates in this work match well with Jiang et al. (2017) in North America and Africa
353 but are 38%, 157% and 228% higher than those in Jiang et al. (2017) in Asia, South America
354 and Australia, respectively. A major discrepancy between this work and Jiang et al. (2017) is
355 the treatment of ocean grids. Jiang et al. (2017) defined ocean grids as continental boundary



356 conditions, which were rewritten hourly using the optimized CO concentrations archived from
357 the suboptimal sequential Kalman filter by assimilating MOPITT CO observations. Only
358 MOPITT data over land were assimilated in the 4D-var assimilations in Jiang et al. (2017)
359 without any change in CO distribution over the ocean. In addition, the large differences in
360 chemical sources and sinks between GC-Adjoint-HEMCO and GC-Adjoint-STD, for example,
361 lower VOC-generated CO emissions by 40-60% and higher CO sinks by 20-40% in GC-
362 Adjoint-HEMCO as shown in Table 2, may also contribute to the discrepancy in the derived a
363 posteriori CO emission estimates.

364 As shown in Fig. 8D, the a posteriori simulation demonstrates positive biases in CO
365 columns over China and Southeast Asia, which is a signal of overestimated local CO emissions;
366 meanwhile, the negative biases over the northern Pacific Ocean are reduced in the a posteriori
367 simulation. The negative biases over the remote ocean are more affected by CO chemical
368 sources and sinks; however, biases in chemical sources cannot be effectively adjusted because
369 of the global uniform scaling factor for CH₄-generated CO emissions; biases in chemical sinks
370 cannot be adjusted because of the fixed OH fields in the tagged-CO simulation. Jiang et al.
371 (2017) tried to address this problem by defining continental boundary conditions so that the
372 inverse analysis is dominated by local MOPITT observations to avoid the influence of model
373 biases accumulated within the long-range transport. Conversely, CO emissions over China and
374 Southeast Asia are overestimated in this work to offset the negative biases over the northern
375 Pacific Ocean. We expect similar overestimations in the a posteriori CO emission estimates
376 over South America, southern Africa, and Australia in this work because it is the effective
377 pathway to reduce the negative bias over the ocean in the Southern Hemisphere.

378 **4. Conclusion**

379 An updated version (GC-Adjoint-HEMCO) of the adjoint of GEOS-Chem model was
380 developed in this work. The major updates of GC-Adjoint-HEMCO include 1) support for the



381 MERRA-2 meteorological data so that we can perform long-term inverse analysis with
382 consistent meteorological data in 1979-present and 2) support for the HEMCO including
383 emission inventories CEDS, MIX, NEI2011, DICE_AF, AF_EDGAR43, APEI and GFED4.
384 The updated emission inventories are critical for reliable inverse analysis, as they are helpful
385 for better convergence of 4D-var assimilations by setting the more reasonable a priori penalty
386 in the cost function. CO emissions from various inventories, chemical sources and sinks, and
387 surface and column CO concentrations provided by GC-Adjoint-HEMCO are compared with
388 those provided by GC-v12 and GC-Adjoint-STD, and we find good agreement in the forward
389 simulations between GC-Adjoint-HEMCO and GC-v12.

390 Two types of OSSE are employed to evaluate the performance of GC-Adjoint-HEMCO
391 in 4D-var assimilations. The a posteriori CO emissions converge to the true states in all major
392 emission regions with fully covered pseudo-CO observations; the limited coverage of
393 observations by sampling the pseudo-CO observations at the locations/times of MOPITT CO
394 observations and smoothing with MOPITT averaging kernels resulted in approximately 15%
395 underestimations in the a posteriori CO emissions over North America and Europe.
396 Furthermore, as an example application of GC-Adjoint-HEMCO, we constrain anthropogenic
397 CO emissions in 2015 by assimilating MOPITT CO observations. The a posteriori
398 anthropogenic CO emission estimates derived in this work match well with Jiang et al. (2017)
399 in North America and Africa but are overestimated in Asia, South America and Australia,
400 which could be associated with the different treatment of MOPITT CO observations over ocean
401 grids and the large differences in CO chemical sources and sinks. The GC-Adjoint-HEMCO
402 developed in this work is a useful extension for the adjoint of GEOS-Chem model. More efforts
403 are needed to support tagged-CO simulation with higher spatial resolutions and to support full
404 chemistry simulation in GC-Adjoint-HEMCO.

405



406 **Code and data availability:** The MOPITT CO data can be downloaded from
407 <https://asdc.larc.nasa.gov/data/MOPITT/>. The GEOS-Chem model (version 12.8.1) can be
408 downloaded from http://wiki.seas.harvard.edu/geos-chem/index.php/GEOS-Chem_12#12.8.1.
409 The adjoint of GEOS-Chem model (GC-Adjoint-STD) can be downloaded from
410 http://wiki.seas.harvard.edu/geos-chem/index.php/GEOS-Chem_Adjoint. The adjoint of
411 GEOS-Chem model (GC-Adjoint-HEMCO) can be downloaded from
412 <https://doi.org/10.5281/zenodo.7512111>.

413

414 **Author Contributions:** Z.J. designed the research. Z.T. developed the model code and
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417

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419

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425

426 **Tables and Figures**

427 **Table 1.** CO emissions for each inventory in 2015 with unit Tg/y.

428

429 **Table 2.** Regional combustion CO emissions, VOC-generated CO (PCO_NMVOC), CH₄-
430 generated CO (PCO_CH₄), CO sinks (CO_{OH}, calculated as CO_{OH} = KRATE×CO×OH),
431 and simulated surface and column CO concentrations in 2015. The region definitions are shown
432 in Fig. 2A.



433

434 **Table 3.** Annual scaling factors of anthropogenic CO emissions in OSSEs. The scaling factors
435 represent the ratio of the estimated to true emissions. The ratio for the first guess is 0.5. The
436 actual value is 1.0. The pseudo-observations are produced by GC-Adjoint-HEMCO forward
437 simulation. The full modeled CO fields are used in OSSE-FullOBS as pseudo-CO observations.
438 The modeled CO fields are smoothed with MOPITT averaging kernels to produce MOPITT-
439 like pseudo-CO observations in OSSE-MOPITT.

440

441 **Table 4.** Regional anthropogenic CO emissions (with unit Tg/y) and annual scaling factors in
442 2015 in this work and Jiang et al. 2017.

443

444 **Fig. 1.** Framework to read the updated emission inventories in GC-Adjoint-HEMCO.

445

446 **Fig. 2.** Total combustion CO emissions in 2015 from (a) GC-v12; (b) GC-Adjoint-HEMCO;
447 (c) GC-Adjoint-STD. The unit is molec/cm²/s.

448

449 **Fig. 3.** Monthly variation in combustion CO emissions in 2015 from GC-v12, GC-Adjoint-
450 HEMCO and GC-Adjoint-STD.

451

452 **Fig. 4.** Averages of surface CO concentrations in 2015 from (a) GC-Adjoint-HEMCO driven
453 by MERRA-2, (b) GC-Adjoint-HEMCO driven by GEOS-FP and (c) their difference; (d-f)
454 same as panels a-c, but for CO columns.

455

456 **Fig. 5.** Averages of surface CO concentrations in 2015 from (a) GC-v12; (b) GC-Adjoint-
457 HEMCO; (c) GC-Adjoint-STD. (d-f) same as panels a-c, but for CO columns.

458

459 **Fig. 6.** (a) Annual scaling factors in OSSE-FullOBS. The scaling factors represent the ratio of
460 the estimated to true emissions. The ratio for the first guess is 0.5. The actual value is 1.0. (b-
461 c) the a priori and a posteriori biases calculated by (model-observation)/observation in OSSE-
462 Full. (d-f) same as panels a-c, but for OSSE-MOPITT.

463

464 **Fig. 7.** (a) Biases in monthly initial CO conditions in 2015 in the original GEOS-Chem
465 simulation. (b) same as panel a, but with optimized initial CO conditions provided by
466 suboptimal sequential Kalman filter. The biases are calculated by (model-MOPITT)/MOPITT.



467

468 **Fig. 8.** (a) A priori anthropogenic CO emissions in 2015 with unit molec/cm²/s; (b) Annual
469 scaling factors for CO emissions in 2015. The scaling factors represent the ratio of the estimated
470 to true emissions. (c-d) the a priori and a posteriori biases calculated by (model-
471 MOPITT)/MOPITT.

472

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Inventories	GC-v12	GC-Adjoint-HEMCO	Inventories	GC-Adjoint-STD
CEDS	613.57	613.85	GEIA	445.88
MIX	321.18	321.71	INTEX-B	353.03
NEI2011	35.83	37.70	NEI2008	52.87
DICE_AF + AF_EDGAR43	83.42	83.02	\	\
APEI	6.10	6.17	CAC	10.20
GFED4	437.13	435.89	GFED3	382.04

Table 1. CO emissions for each inventory in 2015 with unit Tg/y.

Version Region	Combustion Emission (Tg/y)			PCO_NMVOC (kg/s)			PCO_CH4 (kg/s)		
	GC-v12	GC-Adjoint-HEMCO	GC-Adjoint-STD	GC-v12	GC-Adjoint-HEMCO	GC-Adjoint-STD	GC-v12	GC-Adjoint-HEMCO	GC-Adjoint-STD
Asia	320.66	320.38	331.65	15.49	15.52	22.37	14.21	14.40	10.67
North America	73.96	66.93	60.65	7.05	6.83	14.75	7.45	7.66	5.23
Africa	199.51	193.29	179.22	34.57	33.92	52.38	19.57	19.85	16.18
South America	79.04	78.91	75.82	44.15	42.55	74.64	17.14	17.42	14.08
Europe	31.58	30.96	48.48	4.20	4.14	10.17	7.13	7.41	4.58
Australia	12.24	11.99	22.87	21.23	20.68	48.89	13.88	14.62	10.67
Version Region	CO_OH (kg/s)			CO (surface ppbv)			CO (column xco)		
	GC-v12	GC-Adjoint-HEMCO	GC-Adjoint-STD	GC-v12	GC-Adjoint-HEMCO	GC-Adjoint-STD	GC-v12	GC-Adjoint-HEMCO	GC-Adjoint-STD
Asia	52.26	51.34	40.87	179.56	184.29	187.90	90.23	88.16	89.58
North America	23.02	22.57	16.20	120.38	113.49	108.27	79.16	76.27	71.35
Africa	63.78	61.84	51.03	133.56	127.38	141.97	84.26	81.52	86.36
South America	49.06	48.85	41.25	107.98	106.16	132.24	72.93	70.67	78.75
Europe	20.65	20.92	14.27	112.88	111.33	120.09	74.83	72.34	70.45
Australia	31.42	31.98	25.27	67.45	65.00	84.80	56.35	54.02	61.15

Table 2. Regional combustion CO emissions, VOC-generated CO (PCO_NMVOC), CH₄-generated CO (PCO_CH₄), CO sinks (CO_OH, calculated as CO_OH = KRATE×CO×OH), and simulated surface and column CO concentrations in 2015. The region definitions are shown in Fig. 2A.



	Scaling Factors OSSE-FullOBS	Scaling Factors OSSE-MOPITT
Asia	1.00	1.04
North America	0.97	0.88
Africa	0.97	1.01
South America	1.00	1.02
Europe	0.98	0.84
Australia	0.94	0.81

Table 3. Annual scaling factors of anthropogenic CO emissions in OSSEs. The scaling factors represent the ratio of the estimated to true emissions. The ratio for the first guess is 0.5. The actual value is 1.0. The pseudo-observations are produced by GC-Adjoint-HEMCO forward simulation. The full modeled CO fields are used in OSSE-FullOBS as pseudo-CO observations. The modeled CO fields are smoothed with MOPITT averaging kernels to produce MOPITT-like pseudo-CO observations in OSSE-MOPITT.

		Asia	North America	Africa	South America	Europe	Australia
This work	A priori CO emissions	243.53	34.42	23.24	30.39	25.94	2.02
	A posteriori CO emissions	283.20	50.47	35.34	42.92	41.62	2.79
	Scaling Factors	1.16	1.47	1.52	1.41	1.60	1.38
Jiang et al. 2017	A priori CO emissions	270.50	43.70	29.39	17.47	44.45	0.83
	A posteriori CO emissions	205.40	47.06	35.04	16.67	53.58	0.82
	Scaling Factors	0.76	1.08	1.19	0.95	1.21	0.99

Table 4. Regional anthropogenic CO emissions (with unit Tg/y) and annual scaling factors in 2015 in this work and Jiang et al. 2017.

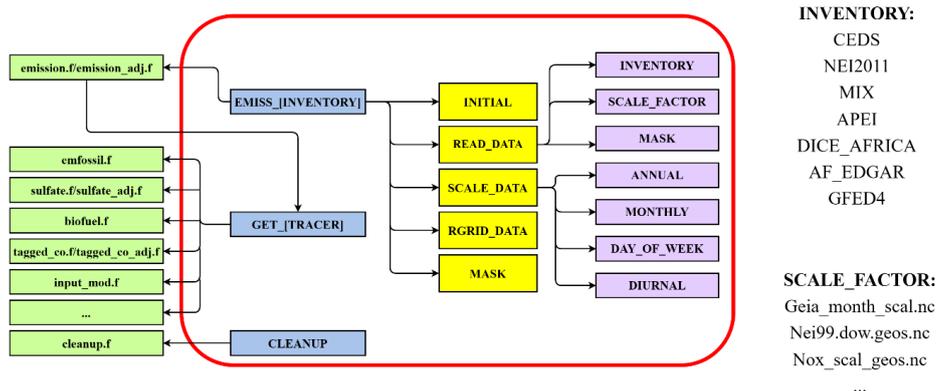


Fig. 1. Framework to read the updated emission inventories in GC-Adjoint-HEMCO.

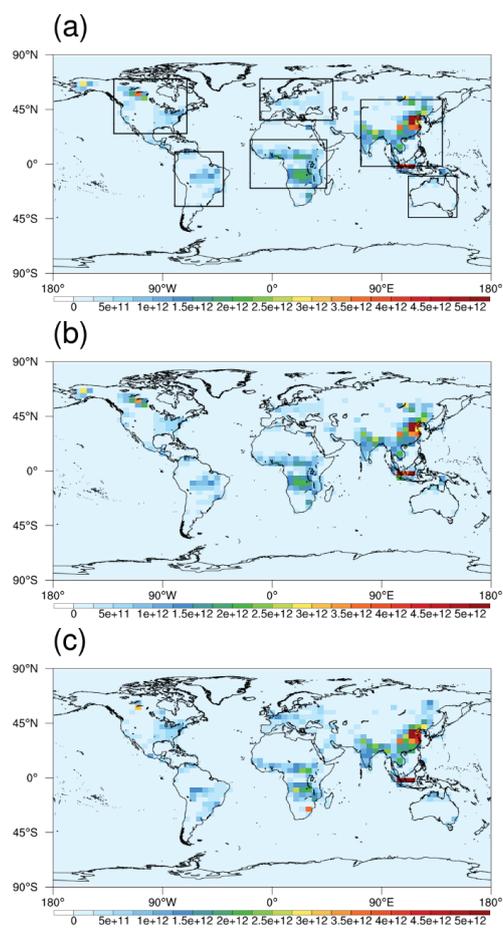


Fig. 2. Total combustion CO emissions in 2015 from (a) GC-v12; (b) GC-Adjoint-HEMCO; (c) GC-Adjoint-STD. The unit is molec/cm²/s.

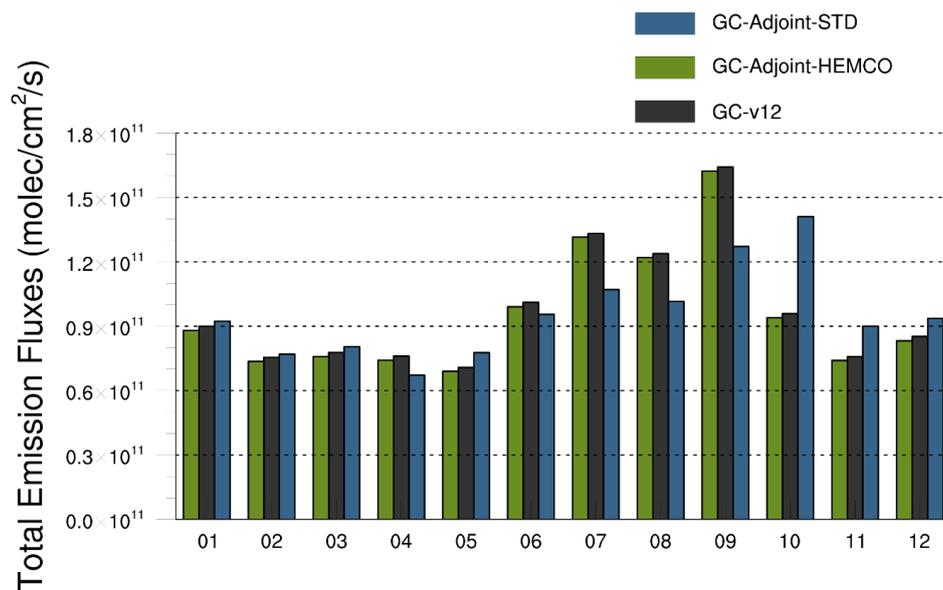


Fig. 3. Monthly variation in combustion CO emissions in 2015 from GC-v12, GC-Adjoint-HEMCO and GC-Adjoint-STD.

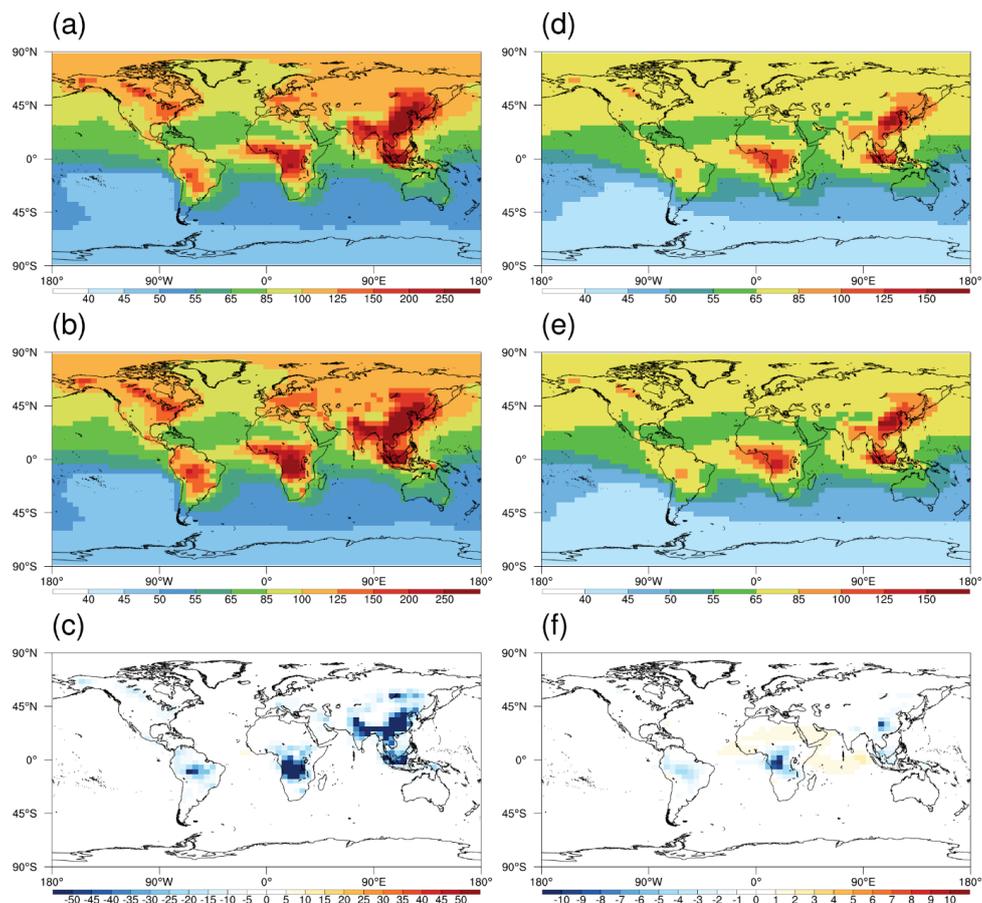


Fig. 4. Averages of surface CO concentrations in 2015 from (a) GC-Adjoint-HEMCO driven by MERRA-2, (b) GC-Adjoint-HEMCO driven by GEOS-FP and (c) their difference; (d-f) same as panels a-c, but for CO columns.

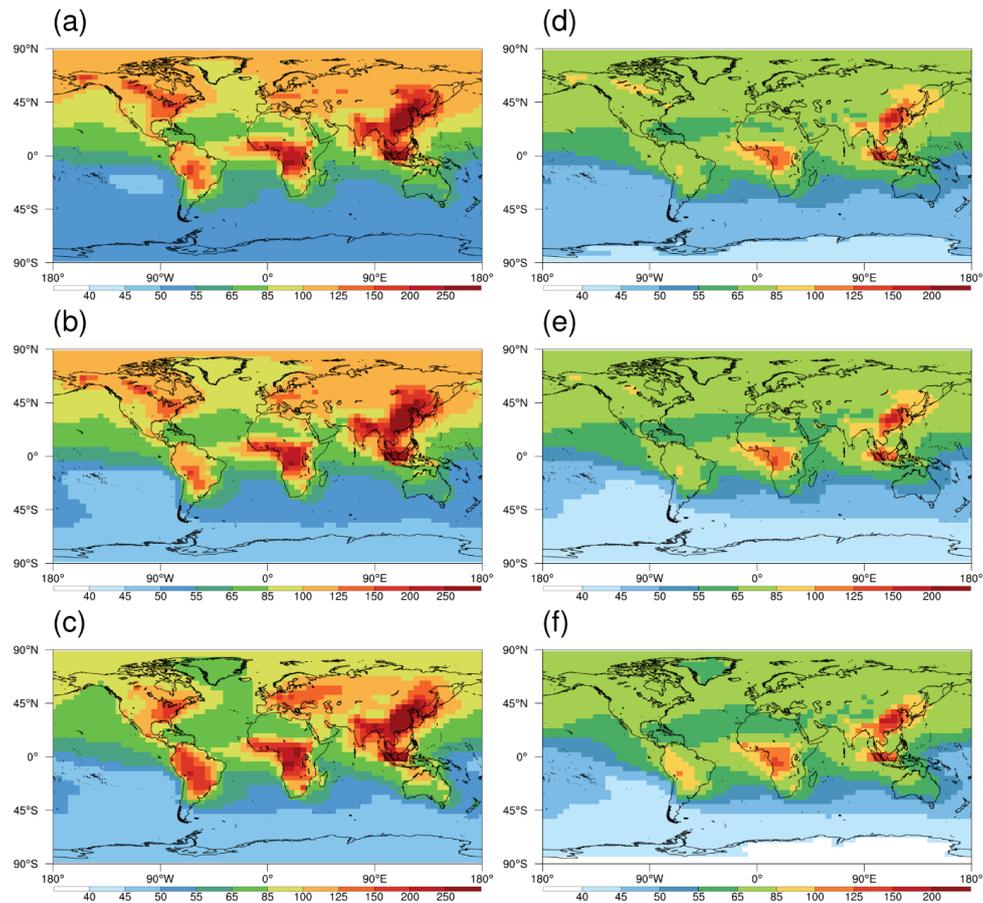


Fig. 5. Averages of surface CO concentrations in 2015 from (a) GC-v12; (b) GC-Adjoint-HEMCO; (c) GC-Adjoint-STD. (d-f) same as panels a-c, but for CO columns.

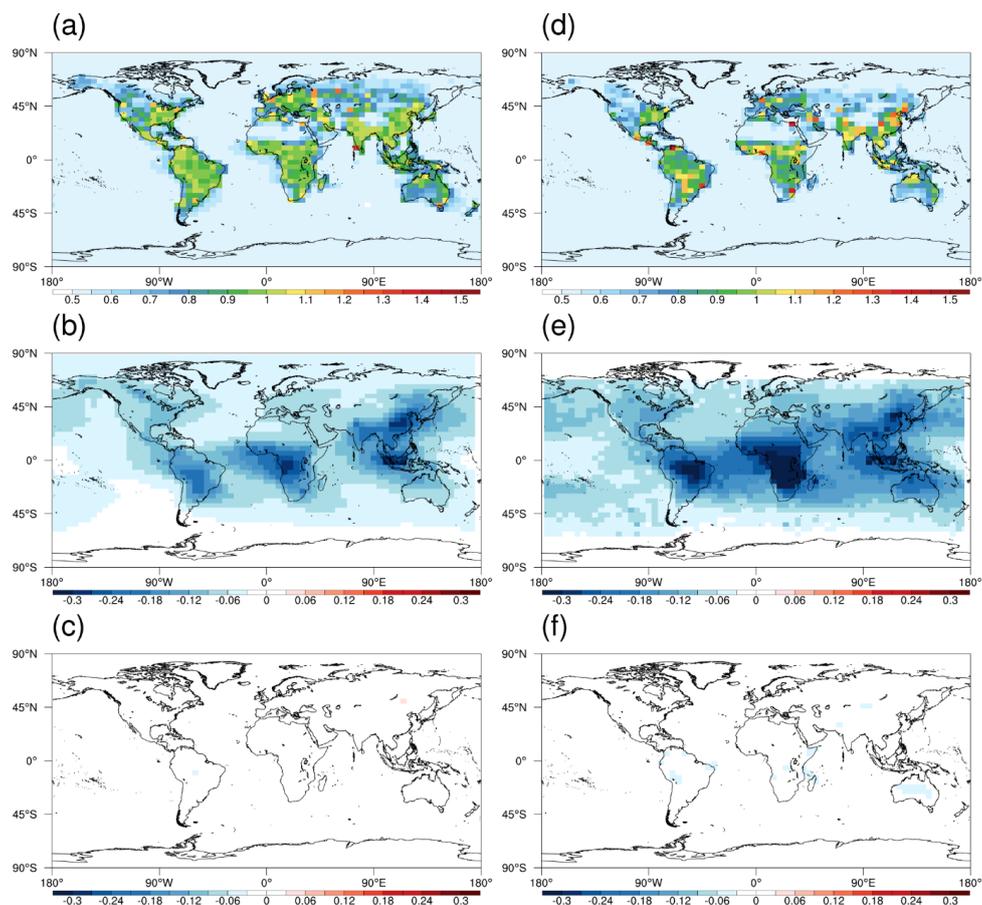


Fig. 6. (a) Annual scaling factors in OSSE-FullOBS. The scaling factors represent the ratio of the estimated to true emissions. The ratio for the first guess is 0.5. The actual value is 1.0. (b-c) the a priori and a posteriori biases calculated by (model-observation)/observation in OSSE-Full. (d-f) same as panels a-c, but for OSSE-MOPITT.

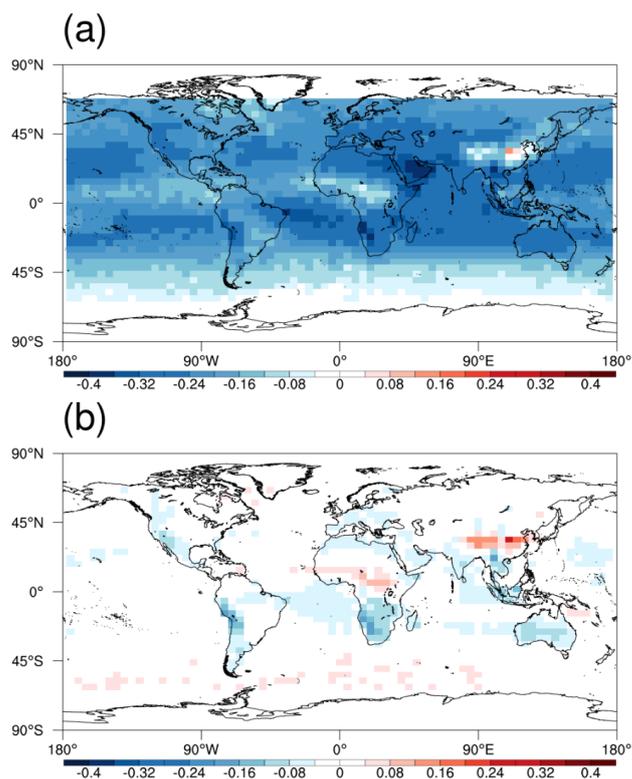


Fig. 7. (a) Biases in monthly initial CO conditions in 2015 in the original GEOS-Chem simulation. (b) same as panel a, but with optimized initial CO conditions provided by suboptimal sequential Kalman filter. The biases are calculated by $(\text{model-MOPITT})/\text{MOPITT}$.

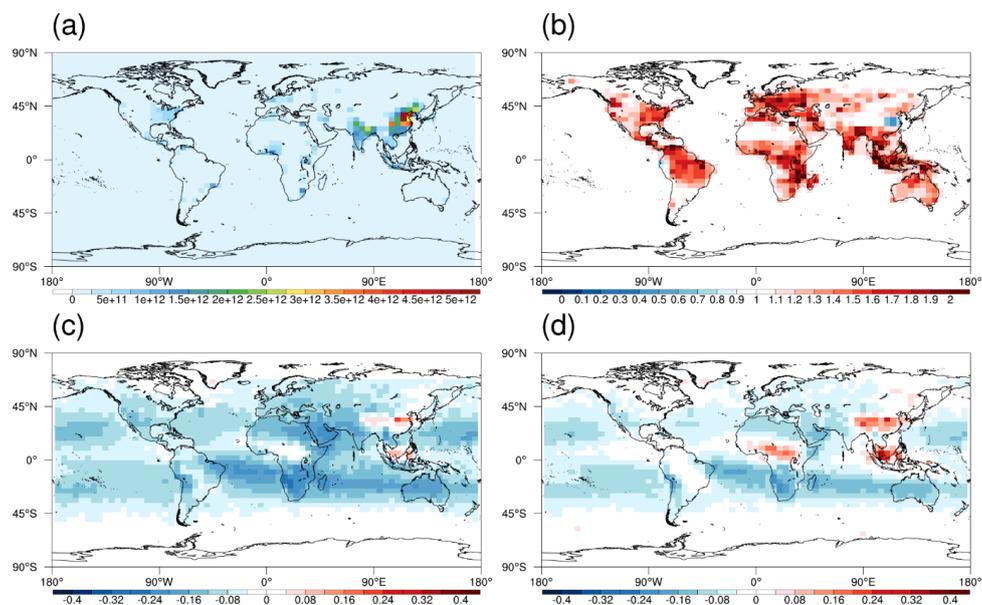


Fig. 8. (a) A priori anthropogenic CO emissions in 2015 with unit molec/cm²/s; (b) Annual scaling factors for CO emissions in 2015. The scaling factors represent the ratio of the estimated to true emissions. (c-d) the a priori and a posteriori biases calculated by (model-MOPITT)/MOPITT.