



Generalized spatiotemporally-decoupled framework for reconstructing the source of non-constant atmospheric radionuclide releases

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9 Abstract. Determining the source location and release rate are critical in assessing the environmental consequences of 10 atmospheric radionuclide releases, but remain challenging because of the huge multi-dimensional solution space. We propose 11 a generalized spatiotemporally-decoupled two-step framework to reduce the dimension of the solution space in each step and 12 improves the reconstruction accuracy, which is applicable to non-constant releases. The decoupling process is conducted by 13 applying a temporal sliding-window average filter to the observations, thereby reducing the influence of temporal variations 14 in the release rate and ensuring that the features of the filtered data are dominated by the source location. A machine learning 15 model is trained to link these features to the source location, enabling independent source localization. Then the release rate is 16 determined using projected alternating minimization with the L1-norm and total variation regularization algorithm. Validation using SCK-CEN ⁴¹Ar experimental data demonstrates that the localization error is less than 1%, and the temporal variations, 17 18 peak release rate, and total release are reconstructed accurately. The proposed method exhibits higher accuracy and a smaller 19 uncertainty range than the correlation-based and Bayesian methods. Furthermore, it achieves stable performance with different 20 hyperparameters and produces low error levels even with only a single observation site.

21 1. Introduction

Atmospheric radionuclide release is a major environmental concern of nuclear industry, including nuclear energy and its heat applications, isotope production and post-processing of radioactive waste. These releases have been observed in the Chernobyl nuclear accident (Anspaugh et al., 1988), and Fukushima nuclear explosion (Katata et al., 2012) with partially-known source information, i.e. the location. Recently, there have been several atmospheric radionuclide leaks from unknown sources, such as the 2017 ¹⁰⁶Ru leakage (Masson et al., 2019) and 2020 ^{134/137}Cs detection in northern Europe (Ingremeau and Saunier, 2022), which have raised global concerns regarding the subsequent hazard to public health. Identification of source information in these events is critical for safe operation of nuclear facilities, consequence assessment, and emergency response.





29 During these events, source information could not be directly measured or determined because of the lack of information 30 on the source of the leak. Instead, source information could only be reconstructed through inversion methods that identified 31 the optimal solution by comparing the environmental observations with atmospheric dispersion simulations using different 32 estimates of the source location and release rate. Such reconstruction simultaneously identifies the source location and release 33 rate because the observations are intuitively determined by both parameters. In this case, the reconstruction searches for the 34 solution over a tremendous multi-dimensional space, where the dimension is the sum of the number of space coordinates and 35 the length of the estimated release window. Therefore, the inversion is weakly constrained and can become ill-posed in the case of spatiotemporally limited observations and uncertainties in the atmospheric dispersion models. Unfortunately, this is 36 37 quite often the case for atmospheric radionuclide releases.

38 To reduce the problem of ill-posedness, most previous studies have attempted to constrain the reconstruction by imposing 39 assumptions on a certain feature of the solution, including both statistical and deterministic assumptions. Statistical assumptions are widely used in Bayesian methods to simultaneously reconstruct the posterior distributions of spatiotemporal 40 41 source parameters (De Meutter et al., 2021; Meutter and Hoffman, 2020; Xue et al., 2017a). This assumes that the model-42 observation discrepancies follow a certain statistical distribution (i.e. the likelihood of Bayesian methods), with the normal 43 (Eslinger and Schrom, 2016; Guo et al., 2009; Keats et al., 2007, 2010; Rajaona et al., 2015; Xue et al., 2017a, b; Yee, 2017; 44 Yee et al., 2008; Zhao et al., 2021) and log-normal (Chow et al., 2008; Dumont Le Brazidec et al., 2020; KIM et al., 2011; 45 Monache et al., 2008; Saunier et al., 2019; Senocak, 2010; Senocak et al., 2008) distributions being two popular choices. Other 46 candidate distributions include Cauchy, log-Cauchy, and T3-10, which have been compared with normal and log-normal 47 distributions in reconstructing the source parameters of the Prairie Grass field experiment (Wang et al., 2017). The results 48 demonstrate that the likelihoods are sensitive to both the dataset and the targeted source parameters. Some studies have 49 constructed the likelihood based on multiple metrics that measure the model-observation discrepancies in an attempt to better constrain the solution (Lucas et al., 2017; Jensen et al., 2019). More sophisticated methods involve the use of different 50 51 statistical distributions for the likelihoods of non-detections and detections (De Meutter et al., 2021; Meutter and Hoffman, 52 2020). Recent studies have suggested the use of log-based distributions and tailored parameterization of the covariance matrix 53 as a means of better quantifying the uncertainties in the reconstruction (Dumont Le Brazidec et al., 2021). These Bayesian 54 methods have been applied to real atmospheric radionuclide releases, such as the 2017 106Ru event, and have provided 55 important insights into the source and release process (Dumont Le Brazidec et al., 2020; Saunier et al., 2019; Dumont Le 56 Brazidec et al., 2021; De Meutter et al., 2021). However, these studies have also revealed that the likelihood in Bayesian 57 methods must be exquisitely designed and parameterized to achieve satisfactory spatiotemporal source reconstruction (Dumont 58 Le Brazidec et al., 2021; Wang et al., 2017). With suboptimal design, the reconstruction may exhibit a bimodal posterior 59 distribution (Meutter and Hoffman, 2020), which remains a challenge for robust applications in different scenarios.

Deterministic assumptions mainly involve entropy (Krysta and Bocquet, 2007; Bocquet, 2005b, a) and a constant release rate (Kovalets et al., 2020, 2018; Efthimiou et al., 2018, 2017; Tomas et al., 2021; Andronopoulos and Kovalets, 2021; Ma et al., 2018). Compared with entropy, the constant-release assumption is more popular and is embedded in many Bayesian





63 methods (Yee et al., 2008; Eslinger and Schrom, 2016; Meutter and Hoffman, 2020; Zhao et al., 2021; De Meutter et al., 2021), 64 substantially reducing the dimension of the solution space to 5 or 6 (i.e. the two or three source location coordinates, the start 65 and end time of the release, and the total release). Recently, the constant-release assumption has been found to enable separate 66 reconstruction of the source location and release rate because the relative spatiotemporal distribution of the simulated 67 atmospheric concentrations is independent of the release rate and is determined by the source location and meteorology 68 (Effhimiou et al., 2017; Kovalets et al., 2018). Because the meteorology can be independently calculated, it is possible to 69 efficiently identify the source location without knowing the specific release rate by simply comparing the correlation 70 coefficients between the observations and simulations using different source locations. The constant release rate is then 71 obtained by calculating a scale factor between the simulation using the identified source location and the observations. This 72 method exhibits high accuracy for constant releases under stationary meteorological conditions, such as synthetic simulation 73 experiments (Ma et al., 2018) and wind tunnel experiments (Kovalets et al., 2018; Efthimiou et al., 2017). However, real 74 releases usually exhibit temporal variations and may experience non-stationary meteorological fields, both of which can lead 75 to noticeable uncertainties in the source localization results of the correlation-based method (Tomas et al., 2021; 76 Andronopoulos and Kovalets, 2021), thus limiting its practical usage. Despite this restriction of the constant-release 77 assumption, the correlation-based method does efficiently reduce the solution space for the source localization and release rate 78 estimation, suggesting the potential for reliable source reconstruction.

79 In this study, we relax the constant-release assumption and propose a more general spatiotemporally decoupled source 80 reconstruction method for non-constant release scenarios. Our approach uses the simple facts that the source location is fixed 81 during the atmospheric radionuclide release process and the peak amplitude of temporal observations is mainly affected by the 82 release rate of a spatially fixed source (Li et al., 2019b). Based on these facts, the proposed method removes the influence of 83 the release rate on the observations through a temporal sliding-window average filter, which approximates the constant-release 84 case and enables decoupled source localization in the absence of release rate estimation. Because the peak amplitude is reduced, 85 existing methods based on direct observation-simulation comparisons may be unable to localize the source. Thus, the response 86 features of the filtered observations are extracted and mapped to the source location by training a data-driven machine learning 87 model using the extreme gradient boosting (XGBoost) algorithm (Chen and Guestrin, 2016). To fully capture the response 88 features at each observation site, tailored time- and frequency-domain features are designed and optimized using the feature 89 selection technique of XGBoost. Using this optimized model, the source is localized based on the filtered observations. Once 90 the source location has been retrieved, the non-constant release rate is determined using the Projected Alternating MInimization 91 with L1-norm and Total variation regularization (PAMILT) algorithm (Fang et al., 2022), which is robust to model 92 uncertainties. The sequential spatiotemporal reconstruction reduces the dimension of the solution space at each step, which 93 helps to improve the accuracy and reliability of the reconstruction.

The performance of the proposed method is compared with the correlation-based method (constant-release assumption) for source localization and the Bayesian method (statistical assumption) for spatiotemporal accuracy. The sensitivity of the source





localization to the spatial search range, size of the sliding window, feature type, and number and combination of sites is also
 investigated.

98 2. Material and Methods

99 2.1 Source reconstruction models

100 For an atmospheric radionuclide release, Eq. (1) relates the observations at each observation site to the source parameters:

101
$$\boldsymbol{\mu} = \mathbf{F}(\mathbf{r}, \mathbf{q}) + \boldsymbol{\varepsilon}, \tag{1}$$

where $\mathbf{\mu} = [\mu_1, \mu_2, \dots, \mu_N] \in \mathbb{R}^N$ is an observation vector composed of observations at *N* sequential time steps, the function **F** maps the source parameters to the observations, i.e. an atmospheric dispersion model, **r** refers to the source location, $\mathbf{q} \in \mathbb{R}^N$ is the temporally varying release rate, and $\mathbf{\varepsilon} \in \mathbb{R}^N$ is a vector containing both model and measurement errors.

In most source reconstruction models, F is simplified to the product of q and a source–receptor matrix A that depends onthe source location:

107
$$\boldsymbol{\mu} = \mathbf{A}(\mathbf{r})\mathbf{q} + \boldsymbol{\varepsilon}$$
, (2)

where $\mathbf{A}(\mathbf{r}) = [A_1(\mathbf{r}), A_2(\mathbf{r}), \dots, A_N(\mathbf{r})]^T \in \mathbb{R}^{N \times N}$ and each row describes the sensitivity of an observation to the release rate q given the source location \mathbf{r} .

110 2.2 Spatiotemporal decoupling

111 A straightforward way to solve Eq. (2) is to simultaneously retrieve the source location and release rate, which the solution 112 space is huge and difficult to constrain. Several studies have pointed out that the source location can be retrieved separately 113 without knowledge of the exact release rate, on the condition that the release rate is constant (Efthimiou et al., 2018; Kovalets 114 et al., 2018; Effhimiou et al., 2017; Ma et al., 2018). The key reason is that, in constant-release cases, the relative spatiotemporal 115 distribution of radionuclides is determined by the meteorological conditions and the relative positions between the source and 116 receptors, and the constant release rate only changes the absolute values. Although the release rate may counteract the influence 117 of the meteorological conditions and relative position at a single observation site, it cannot change the whole spatiotemporal 118 distribution at multiple observation sites. Therefore, by analysing the spatiotemporal distribution of radionuclides at multiple 119 observation sites, it is possible to localize the source without knowing the release rate under the constant-release assumption. 120 To provide a more general method, we take advantage of the fact that the source location has been fixed during all known 121 atmospheric radionuclide releases, such as the Chernobyl nuclear accident (Anspaugh et al., 1988), the Fukushima nuclear

- 122 explosion (Katata et al., 2012), and the 2017¹⁰⁶Ru leakage (Masson et al., 2019). With a fixed source location, the release rate
- mainly influences the peak values of the temporal observations at each site, whereas the meteorology determines the timing of
- 124 the peaks (Li et al., 2019b). Therefore, the constant-release case can be approximated by reducing the influence of the release





rate, enabling separate source localization and release rate estimation in a more general case and reducing the solution space at each step. For this purpose, we introduce an operator matrix $\mathbf{P} \in \mathbb{R}^{N \times N}$ to reduce the temporal variations of $\mathbf{A}(\mathbf{r})\mathbf{q}$:

127
$$\boldsymbol{\mu}_{p} = \mathbf{P}\boldsymbol{\mu} = \mathbf{P}\mathbf{A}(\mathbf{r})\mathbf{q} + \mathbf{P}\boldsymbol{\varepsilon},$$

(3)

where μ_p refers to the decoupled observations. In this study, the following operator matrix is constructed to impose a temporal sliding-window average filter (Eamonn Keogh, Selina Chu, 2004):

$$130 \quad \mathbf{P} = \frac{1}{T} \begin{bmatrix} 1 & & & & & \\ 1 & 1 & & & & \\ 1 & 1 & \cdots & 1 & & \\ 1 & 1 & \cdots & 1 & 1 & & \\ & 1 & 1 & \cdots & 1 & 1 & & \\ & & 1 & 1 & \cdots & 1 & 1 & \\ & & & \ddots & \ddots & \ddots & \ddots & \\ & & & & 1 & 1 & \cdots & 1 & 1 \\ & & & & & 1 & 1 & 1 & 1 \end{bmatrix},$$

$$(4)$$

where *T* is the size of the sliding window. Although a sliding-window average filter is used in this study, Eq. (3) is compatible with more advanced processing methods, thereby providing a general framework for the spatiotemporal decoupling of μ .

133 **2.3 Source localization without knowing the exact release rates**

134 After applying the filter in Eq. (4), the local peak amplitude of the observations is smoothed out, but the influences of the 135 source position and meteorology remain relatively unchanged, as they determine the long-term temporal trends of observations 136 and are less affected by the filter. The meteorology is known, so it becomes possible to localize the source using the filtered 137 observations. Nevertheless, the specificity of those localization methods that rely on direct observation-simulation 138 comparisons may be substantially compromised because the peak amplitude is reduced. A better choice for source localization 139 would be to use the response features of the filtered observations, which preserve most of the location information. Therefore, 140 it is necessary to establish a link between the response features of the filtered observations and the source location. To achieve 141 this, we train an XGBoost model that maps the response features of the filtered observations to the coordinates of the source.

142 XGBoost is an optimized distributed gradient boosting library. Suppose $D = \{(\mathbf{X}_i, \mathbf{r}_i)\}(|D| = n, \mathbf{X}_i \in \mathbb{R}^p, \mathbf{r}_i \in \mathbb{R}^2)$, where 143 the number of samples is *n* and each sample contains *p* features. \mathbf{X}_i is the given input feature vector of the *i*-th sample and 144 $\mathbf{r}_i = (x_i, y_i)$ is the location vector. XGBoost typically uses multiple decision trees (Fig. 1) to fit the target, which can be 145 formulated as:

146
$$\hat{\mathbf{r}}_i = G(\mathbf{X}) = \sum_{k=1}^K f_k(\mathbf{X}_i), f_k \in \mathcal{F},$$
 (5)

where *K* is the number of trees, $\mathcal{F} = \{f(x) = \omega_{Q(x)}\}(Q: \mathbb{R}^p \to M, \omega \in \mathbb{R}^M)$ is the space of decision trees, and *Q* represents the structure of each tree, mapping the feature vector to *M* leaf nodes. Each f_k corresponds to an independent tree structure *Q*





149 with leaf node weight $\boldsymbol{\omega}$. Equation (5) is then used to predict $\hat{\mathbf{r}}_i = (\hat{x}_i, \hat{y}_i)$ for the *i*-th sample.



150

151 Figure 1. Flowchart of XGBoost for predicting $\hat{\mathbf{r}}_i$ based on decision tree model.

152 XGBoost trains $G(\mathbf{X})$ in Eq. (5) by continuously fitting the residual error until the following objective function is minimized:

153
$$Obj^{(t)} = \sum_{i=1}^{n} \left(\mathbf{r}_{i} - \left(\hat{\mathbf{r}}_{i}^{(t-1)} + f_{t}(\mathbf{X}_{i}) \right) \right)^{2} + \sum_{i=1}^{t} \Omega(f_{i}) , \qquad (6)$$

154 where t represents the training of the t-th tree and $\Omega(f_i)$ is the regularization term, given by:

155
$$\Omega(f) = \Upsilon T + \frac{1}{2}\lambda \sum_{j=1}^{T} \omega_j^2 , \qquad (7)$$

The minimization of Eq. (6) provides the parametric model $G(\mathbf{X})$ that maps the feature ensemble \mathbf{X} extracted from $\boldsymbol{\mu}_p$ to the source location \mathbf{r} .

To comprehensively evaluate the influence of the source location, both time- and frequency-domain features (as outlined in Table 1) are considered during the training process and mapped to the source location by $G(\mathbf{X})$. Among the time-domain features, the wave rate quantifies the amplitude of fluctuations in $\boldsymbol{\mu}_p$, while the temporal mean and median values represent the central moment of $\boldsymbol{\mu}_p$. Additionally, the sample entropy measures the complexity of $\boldsymbol{\mu}_p$, with a lower sample entropy indicating greater self-similarity and less randomness in $\boldsymbol{\mu}_p$. The frequency-domain features are calculated based on the fast Fourier transform (FFT). The FFT mean is the mean value of the Fourier coefficients for $\boldsymbol{\mu}_p$ and the FFT shape mean describes the shape of the Fourier coefficients. These quantities are formulated as follows:





165 FFT mean
$$= \frac{1}{N} \sum_{k=1}^{N} |\mu_{ik}|$$
, (8)
166 FFT shape mean $= \frac{1}{N^{N-1-1}} \sum_{k=1}^{N} k |\mu_{ik}|$, (9)

166 FFT shape mean
$$= \frac{1}{\sum_{k=1}^{N} |\mu_{ik}|} \sum_{k=1}^{N} k |\mu_{ik}|$$
,

where μ_{ik} is the Fourier coefficient and N is the length of μ_p . These features are calculated from the simulated observations at 167

168 each site and provided to XGBoost as initial inputs.

169 Table 1. Summary of the basic information on the observation series features.

Attribute	Feature	Description			
	Wave rate	Difference between 90-th and 10-th quantile of normalized observation series			
Time domain	Mean	Temporal mean value of observation series			
Time domain	Median	Temporal median value of observation series			
	Sample entropy	Complexity of observation series			
Fraguanay domain	FFT mean	Amplitude of power spectral density by FFT			
Frequency domain	FFT shape mean	Shape of power spectral density by FFT			

170 2.4 Release rate estimation

- 171 Once the source location has been retrieved, many existing methods can be used to inversely estimate the release rate. In this
- study, we choose the recent PAMILT method (Fang et al., 2022) because it can correct the intrinsic model errors of the release 172 173 rate estimation and reduce the propagation of localization errors into the release rate estimate.

174 2.5 Numerical implementation

175 2.5.1 Pre-screening of potential source locations

176 To reduce the computational cost and remove low-quality samples, the search range for the source location is pre-screened by 177 evaluating the correlation coefficients between the observations and atmospheric dispersion model simulations, where the 178 source locations are randomly sampled in the considered calculation domain. Because the release rate is unknown, it is assumed 179 to be 1 Bq/h for all simulations. Those source locations corresponding to correlation coefficients above the 40th percentile are 180 selected as the search range of the subsequent refined source localization using XGBoost.

181 2.5.2 Samples for training XGBoost

182 The samples for training $G(\mathbf{X})$ in Eq. (5) are generated based on the simulations described in Sect. 2.5.1, and the source





locations of these simulations are within the search range as determined according to Sect. 2.5.1. The simulation data are scaled by a constant factor (the ratio between the median value of all observations and that of the simulations using a unit release rate), which ensures that the simulations and observations have the same order of magnitude. Gaussian noise is added to the simulation data to simulate the statistical fluctuations of radiation measurements. Those simulations between the first and last data points above the noise level are filtered by a temporal sliding-window average filter with a window size of 5, yielding samples for feature extraction as described in Sect. 2.3.

189 2.5.3 Automatic optimization of XGBoost model

The XGBoost model for source localization is automatically optimized with respect to the hyperparameters and feature selection. Specifically, the Bayesian optimization algorithm is used to optimize the hyperparameters by minimizing the following generalization coefficient (GC) defined under the five-fold cross-validation framework:

193
$$GC = (1 - MCV)^2 + Var(R_k^2),$$
 (10)

194
$$MCV = \frac{1}{5} \sum_{k} R_k^2$$
, (11)

where R_k^2 is the goodness of fit and k is the index of each fold (k = 1, 2, ..., 5). MCV is the mean cross-validation score R_k^2 among the five folds and $Var(R_k^2)$ measures the variance of R_k^2 . This function aims to balance the average and the variance of R_k^2 , thus enhancing the generalization ability of the XGBoost model. In this study, the optimized hyperparameters include max_depth (maximum depth of a tree), learning_rate (step size shrinkage when updating), n_estimators (number of decision trees), min_child_weight (minimum sum of sample weight of a child node), subsample (subsample ratio of the training samples), colsample_bytree (subsample ratio of columns when constructing a tree), reg_lambda (L2 regularization term on weights), and gamma (minimum loss reduction required to split the tree).

The initial input features (Table 1) are optimized by recursive feature elimination with cross-validation (Akhtar et al., 2019), which sorts the features in order of importance and removes the least-important features based on the MCV results. The overall flowchart of the proposed spatiotemporally decoupled source reconstruction model is shown in Fig. S1.

205 2.6 Validation case

206 **2.6.1 SCK-CEN**⁴¹Ar field experiment

The proposed methodology was validated against the observations of the SCK-CEN ⁴¹Ar field experiment, which was carried out at the BR1 research reactor in Mol, Belgium, in October 2001 as a collaboration between NKS and the Belgian Nuclear Research Centre (SCK-CEN) (Rojas-Palma et al., 2004). The major part of the experiment was conducted on 3–4 October, during which time ⁴¹Ar was emitted from a 60-m stack with a release rate of approximately 1.5×10^{11} Bq/h. Meteorological data such as wind speed and direction were provided by the on-site weather mast. For most of the experimental period, the

atmospheric stability was neutral, and the wind was blowing from the southwest. As illustrated in Fig. 2, the source coordinates





213 were (650 m, 210 m).



214

Figure 2. Release location and observation sites of SCK-CEN ⁴¹Ar experiment. The map was created based on the relative positions of the release source and observation sites, as detailed in (Drews et al., 2002). It was plotted using MATLAB 2016b, instead of created by a map provider.

The 60-s-average fluence rates were continuously collected by an array of NaI (Tl) gamma detectors, with different observation sites used on the two days. To convert the measured fluence rates to gamma dose rates (mSv/h), we used the ⁴¹Ar parameters of a previous study (Li et al., 2019a): $E_{\gamma} = 1.2938$ MeV, $f^n(E_{\gamma}) = 0.9921$, $\mu_a = 2.05 \times 10^{-3}$ m⁻¹, and $\omega =$ 7.3516 × 10⁻¹ Sv/Gy. More details of these measurements can be found in reference (Rojas-Palma et al., 2004).

222 **2.6.2** Simulation settings of atmospheric dispersion model

The Risø Mesoscale PUFF (RIMPUFF) model was employed to simulate the dispersion of radionuclides and to calculate the dose rates at each observation site (Thykier-Nielsen et al., 1999). The simulations used on-site measured meteorological data and the modified Karlsruhe–Jülich diffusion coefficients. The calculation domain measured 1800 m×1800 m and the grid resolution was 10 m×10 m. Other RIMPUFF calculation settings followed those of a previous study (Li et al., 2019a), and have been validated against the observations.

228 To establish the datasets for the XGBoost model, 2000 samples and 1000 samples with different source locations were





calculated by RIMPUFF for Oct. 3 and Oct. 4 respectively. The source locations were sampled from the shaded zones in Fig. 2, which were determined according to the positions of the observation sites and the upwind direction. As described in Sect. 2.5.1, we calculated the correlation coefficient for each sample and preserved samples with correlation coefficients greater than the 40th percentile of all correlation coefficients (i.e. 800 samples for Oct. 3 and 400 samples for Oct. 4). The constant factors mentioned in Sect. 2.5.2 are 1.53×10^{11} and 1.48×10^{11} for Oct. 3 and Oct. 4, respectively.

234 2.7 Sensitivity study

235 (1) Search range

The search range is controlled by the pre-screening threshold. The source localization is implemented with pre-screening thresholds determined by the 20th, 40th, 50th, 60th, 80th, and 100th percentiles of the correlation coefficients in the prescreening step, where a lower percentile corresponds to a more refined search range.

239 (2) Size of the sliding window

Temporal filtering with different sliding-window sizes is applied to decouple the source localization from the release rate estimation. In this study, the size of the sliding window ranges from 3 to 10. With these decoupled data, the XGBoost model is trained using the same pattern for the source localization.

243 (3) Feature type

The XGBoost model is trained using only time-domain features and only frequency-domain features, respectively, to investigate the influence of these features on the source localization. The performance of the time-feature-only and frequencyfeature-only models is compared with the all-features result.

247 (4) Number and combination of observation sites

The XGBoost model is trained and applied to the source localization with different numbers of observation sites, namely a single site, two sites, and three sites. For the two- and three-site cases, the model is trained using different combinations of sites and the source location is estimated accordingly.

In all the sensitivity tests, the source location is estimated 50 times with randomly initialized hyperparameters to demonstrate the uncertainty range of the proposed method under different circumstances. The performance of source localization is compared quantitatively using the metrics specified in Sect. 2.8.3.

254 **2.8 Performance evaluation**

255 **2.8.1 Decoupling**

256 The feasibility of decoupling was demonstrated using both the synthetic and real observations of the SCK-CEN ⁴¹Ar field

257 experiment. The former were generated by a simulation using a synthetic temporally varying release profile with sharp increase,

- stable, and gradual decrease phases (as illustrated in Fig. S2), which is typical for an atmospheric radionuclide release (Davoine
- and Bocquet, 2007). The simulations corresponding to the synthetic and real observations should first be processed following





the procedure in Sect. 2.5.2. The decoupling performance is evaluated by comparing the simulation–observation differences before and after the decoupling step. Several statistical metrics can be used to quantify this difference, including the normalized mean square error (NMSE), Pearson's correlation coefficient (PCC), and the fraction of predictions within a factor of 2 and 5 of the observations (FAC 2 and FAC 5, respectively) (Chang and Hanna, 2004).

264 **2.8.2 Optimization of the XGBoost model**

The hyperparameters are optimized with respect to the GC in Eq. (10) and the features are optimized with respect to the MCV in Eq. (11). Larger values of MCV and smaller values of GC indicate better optimization performance. In addition, the importance of each feature to the XGBoost training is evaluated with the built-in *feature importance* measure of the XGBoost model.

269 **2.8.3 Source reconstruction**

270 The relative errors of source localization (δ_r) and total release (δ_o) are calculated to evaluate the source reconstruction accuracy:

271
$$\delta_{\mathbf{r}} = \frac{|\mathbf{r}_{true} - \mathbf{r}_{est}|}{L_D} \times 100\% , \qquad (12)$$

272
$$\delta_Q = \frac{|Q_{true} - Q_{est}|}{Q_{true}} \times 100\%, \qquad (13)$$

where \mathbf{r}_{true} and Q_{true} refer to the real source location and total release of the SCK-CEN ⁴¹Ar field experiment and \mathbf{r}_{est} and Q_{est} are the estimated location and total release, respectively. L_D represents the range of the source domain, which is the distance between the lower and upper borders of the sampled zone, equal to 1034.8 m and 565.7 m on Oct. 3 and Oct. 4, respectively. In addition to the total release, the reconstructed release rates are also compared with the true value in terms of the temporal profile.

278 **2.8.4 Comparison with the Bayesian method**

The proposed method is compared with the popular Bayesian method based on the SCK-CEN ⁴¹Ar experiment, with the same search range used for source localization in both methods (Fig. 2). The Bayesian method is augmented with an in-loop inversion of the release rate at each iteration step of the Markov chain Monte Carlo sampling. The prior distribution of the Bayesian method is a uniform distribution and the likelihood is a log-Cauchy distribution. More detailed information is presented in Supplementary Note S1.

284 **2.8.5 Uncertainty range**

The uncertainty ranges are calculated and compared for the correlation-based method, the Bayesian method, and the proposed method. For the correlation-based method, the uncertainty range is calculated using the source locations with the top-50





correlation coefficients. For the proposed method, the uncertainty range is calculated from 50 Monte Carlo runs with randomly initialized hyperparameters. The Bayesian method itself provides the uncertainty range through the posterior distribution. To be consistent with the other two methods, the results with the top-50 frequencies are selected for the comparison.

290 **3. Results and Discussion**

291 **3.1 Decoupling performance**

Figure S3 displays the original and filtered observations at different observation sites for both days. The results demonstrate that the peak values have been smoothed out and the long-term trends are preserved to a large degree.

Figure 3 compares the decoupling performance for both the synthetic and real observations, by plotting the constant-release simulations against the observations before and after decoupling. For the synthetic observations, the decoupled data are more concentrated along the 1:1 line for both days, and all the decoupled data fall within the 2-fold lines for Oct. 3. For real observations, the dots before decoupling in Fig. 3 have a dispersed distribution for both Oct. 3 and Oct. 4, indicating limited correlations with the simulations. After decoupling, the dots are more concentrated along the 1:1 and 1:2 (2:1) lines. These phenomena indicate a noticeably increased agreement between the decoupled observations and the constant-release simulations.







Figure 3. Scatter plots of the original (yellow squares) and decoupled (green squares) observations versus the constant-release simulation results. (a) Oct. 3-Synthetic observations; (b) Oct. 4-Synthetic observations; (c) Oct. 3-Real observations; (d) Oct. 4-Real observations.

Table 2 quantitatively compares the agreements presented in Fig. 3. For both the synthetic and real observations, all metrics are greatly improved after decoupling, especially NMSE and PCC, confirming the better agreement between the decoupled observations and the constant-release simulations. The decoupling performs better with the synthetic observations than with the real observations. This is because the synthetic observations are free of measurement errors. The improved agreement indicates that the decoupling step significantly reduces the influence of temporal variations in release rates across the real observations.

310 **Table 2.** Quantitative metrics for the decoupling validation.

	Experiment		NMSE	PCC	FAC2	FAC5
Synthetic observations	Oct. 3	Before decoupling	0.69704	0.5315	0.7647	0.8235
		After decoupling	0.0239	0.9514	1	1
	Oct. 4	Before decoupling	0.9290	-0.0267	0.7292	0.7292
		After decoupling	0.0956	0.6179	0.9412	0.9779
Real observations	Oct. 3	Before decoupling	1.4437	0.3572	0.3824	0.5147
		After decoupling	0.2730	0.6976	0.7273	0.8864
	Oct. 4	Before decoupling	1.9290	-0.2099	0.3073	0.4948
		After decoupling	0.3668	0.2802	0.6552	0.9310

311 **3.2 Optimization of XGBoost model**

312 3.2.1 Hyperparameters

- 313 Table S1 summarizes the optimal hyperparameters and corresponding GCs used for source localization in this study. The
- optimal GC on Oct. 3 is smaller than that on Oct. 4, indicating better fitting performance. This difference is possibly the result
- 315 of the larger training dataset for Oct. 3.

316 **3.2.2 Feature selection**

Figure S4 shows the variation of MCV with the number of features for the x and y coordinates. The MCV first increases with the number of features, and then decreases slightly after reaching the maximum for both days. The optimal number of features for Oct. 4 is noticeably smaller than for Oct. 3. In addition, the selected features for Oct. 3 involve all four sites, whereas those for Oct. 4 involve three sites for x and two sites for y. The reduced feature and site numbers indicate a high level of redundancy

321 in the observations acquired on Oct. 4. This is because the observation sites are parallel to the downwind direction and provide





322 similar location information in the crosswind direction.

Figure 4 compares the importance of the selected features at each site. For both days, the temporal features are dominant. For Oct. 3, Site B is the most important, possibly because it is farthest away in the crosswind direction. For Oct. 4, the four sites provide redundant feature information, and many features are removed. This is because the distribution of observation sites is almost parallel to the wind direction on this day. According to Fig. S3, the measurements from Site A and B have a high correlation, thus leading to the removal of features from Site A on Oct. 4. In summary, the feature selection process adapts XGBoost to different application scenarios.











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Figure 4. Feature importance. (a) Oct. 3; (b) Oct. 4.

331 **3.3 Source reconstruction**

332 **3.3.1 Localization**

333 Figure 5 compares the best-estimated source locations of the correlation-based method, the Bayesian method, and the proposed





334 method with the ground truth. The pre-screening zone covers the true source location for both days, but the areas with the 335 highest correlation coefficients are still too large for the point source to be accurately located. The locations with the maximum 336 correlation exhibit errors of 270.19 m and 36.06 m for Oct. 3 and Oct. 4, respectively, indicating that the correlation-based 337 method may produce biased results in the case of non-constant releases. The Bayesian method estimates the location with 338 errors of 19.62 m and 52.81 m for Oct. 3 and Oct. 4, respectively. In comparison, the proposed method achieves the best 339 performance among all the methods. The estimates without feature selection are only 10.65 m (Oct. 3) and 20.62 m (Oct. 4) away from the true locations. Feature selection further reduces these errors to 6.19 m (Oct. 3) and 4.52 m (Oct. 4), which are 340 341 below the grid size (10 m \times 10 m) of the ATDM simulation. The proposed method gives a relative error of less than 0.9% for 342 both days, whereas the Bayesian method produces a relative error of above 11% and that of the correlation-based method can 343 be as high as 26%. The best estimates for Oct. 3 are more accurate than those for Oct. 4, possibly because of the better layout 344 of observation sites (Fig. 2) and the better decoupling results (Fig. 3).



345

Figure 5. Source localization results. The yellow dots denote the maximum correlation points, which are the localization results of the correlation-based method. The green and red stars represent the localization results based on XGBoost before and after feature selection, respectively. The cyan diamonds represent the localization results based on the Bayesian method. (a) Oct. 3; (b) Oct. 4.

349 3.3.2 Release rates

Figure 6 displays the release rates estimated by the Bayesian and PAMILT methods based on the source localization results in Fig. 5. The release rates provided by the Bayesian method present several sharp peaks, corresponding to overestimates of up





to 269.03% (Oct. 3) and 532.35% (Oct. 4). Furthermore, the Bayesian estimates exhibit unrealistic oscillations in the stable release phase. In contrast, the PAMILT method successfully retrieves the peak releases without oscillations for both days. Both the Bayesian and PAMILT estimates give delayed start times of the release, but accurately estimate the end time, especially for Oct. 3. The PAMILT estimate underestimates the total release by 30.01% and 45.95% for Oct. 3 and Oct. 4, respectively; these values are reduced to about 23.83% and 30.60% after feature selection. The Bayesian method gives better total releases because of the overestimated peaks.



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Figure 6. Release rate estimation results with different location estimates. (a) Oct. 3; (b) Oct. 4. The rectangles inside each figure present the location estimates used in the release rate estimations. The green and red stars denote the source locations estimated by XGBoost without and with feature selection, respectively.

362 **3.3.3 Uncertainty range**

363 Figure 7 compares the spatial distribution of 50 estimates produced by the different source localization methods. The estimates 364 of the correlation-based method are highly spread for both days, leading to a highly uniform distribution of the x coordinate 365 for Oct. 3 and two separate distributions of both the coordinates for Oct. 4. The Bayesian method produces a multimodal 366 distribution for both days, in which the estimates are more concentrated than those of the correlation-based method. The 367 corresponding full posteriori distributions in Fig. S5 better reveal the multimodal feature of the Bayesian method, with several 368 peaks of similar probabilities in the estimates of both coordinates on Oct. 3 and the y coordinate on Oct. 4. The multimodal 369 feature indicates the difficulty of constraining the solution in simultaneous spatiotemporal reconstruction, as reported in a 370 previous study (Meutter and Hoffman, 2020). In comparison, the proposed method provides the most concentrated source 371 location estimates. The feature selection moves the centre of the distribution closer to the true location and narrows the 372 distribution of the estimates, especially for Oct. 4.







373



375 Figure 8 compares the uncertainty range and the mean total release of the release rate estimations. For Oct. 3, the Bayesian 376 estimates significantly overestimate the mean values and have a large uncertainty range, whereas the mean PAMILT estimate 377 is very close to the true release and the uncertainty range is smaller than that of the Bayesian method. For Oct. 4, the mean 378 Bayesian estimate exhibits more deviations than the mean PAMILT estimate. Feature selection improves the mean estimate 379 and reduces the uncertainty range of PAMILT because it improves the source localization, thus reducing the deviation in the 380 inverse model of the release rate. On Oct. 3 and Oct. 4, the PAMILT method underestimates the total release by 18.30% and 381 47.42%, respectively, whereas the Bayesian method gives overestimations of 153.61% and 42.29%, respectively. These results 382 demonstrate that the PAMILT method is robust to localization deviations and can reconstruct the timing, peaks, and total 383 release with relatively high accuracy. This robustness reduces the propagation of localization errors to the release rate 384 estimation, and improves the accuracy of spatiotemporally decoupled source reconstruction.







385

Figure 8. Release rate estimates over 50 calculations. (a) Oct. 3; (b) Oct. 4. The shadow represents the release rate range between the minimum and the maximum.

388 Table 3 lists the mean and standard deviation of the relative errors for the 50 estimates given by the various methods. In 389 terms of source localization, the correlation-based method produces the largest mean relative error and standard deviation. The 390 proposed method gives the smallest mean error, about half that of the Bayesian method. Its standard deviation is three-fold 391 smaller than that of the Bayesian method for Oct. 3, but slightly larger for Oct. 4. For the total release, the PAMILT method 392 gives a better standard deviation of the relative error for both days and a better mean relative error for Oct. 3, whereas the Bayesian method produces a better mean relative error for Oct. 4. Feature selection reduces the mean relative error, except for 393 394 the total release for Oct. 3, and slightly increases the standard deviation of the source location and the total release results for 395 Oct. 3.

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- 399





400 **Table 3.** Relative errors of source reconstruction.

Experiments	Statistical parameters (Relative error)		Correlation-based method	Bayesian method	The proposed method		
					XGBoost	XGBoost+ feature selection	
Oct. 3	Source location (δ_r)	Mean	14.10%	11.88%	5.18%	4.68%	
		Std	11.37%	7.53%	1.79%	2.05%	
	Total release (δ_Q)	Mean	-	177.46%	18.03%	18.45%	
		Std	-	167.66%	7.13%	7.68%	
Oct. 4	Source location $(\delta_{\mathbf{r}})$	Mean	14.30%	12.83%	6.83%	4.71%	
		Std	9.60%	1.68%	1.76%	1.53%	
	Total release (δ_Q)	Mean	-	42.29%	54.12%	47.42%	
		Std	-	15.05%	6.47%	5.85%	

401 **3.4 Sensitivity analysis results**

402 **3.4.1 Sensitivity to the search range**

Figure 9 displays the localization error obtained using different pre-screening thresholds to determine the search range. The error is smaller with a lower threshold, implying that a small pre-screening range helps reduce the mean and median errors. As the threshold increases, the mean and median errors, as well as the error range, show an overall tendency to increase, but not in a strictly monotonic way. The mean/median error is less than 12% for Oct. 3 and less than 22% for Oct. 4, indicating robust performance in these tests. Feature selection reduces the mean/median, range, and lower bound of the errors in most tests, demonstrating its efficiency.







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Figure 9. Distribution of relative error (%) over 50 runs with different search ranges. The blue and red solid lines denote average relative error (%) and median relative error (%), respectively. The upper and lower boundaries represent the upper and lower quartiles of relative error (%), respectively. The fences are 1.5 times the inter-quartile ranges of the upper/lower quartiles. The red circles denote data that are not included between the fences. (a) Oct. 3; (b) Oct. 4.

414 **3.4.2** Sensitivity to the size of the sliding window

415 Figure 10 shows the localization errors obtained with different sliding-window sizes. The mean/median error is less than 8% 416 for Oct. 3 and less than 11% for Oct. 4, both of which are smaller than for the various pre-screening ranges. This indicates that 417 the proposed method is more robust to this parameter than to the pre-screening range. For both days, the lowest mean/median 418 and error range occur with relatively large window sizes, i.e. window size of 9 for Oct. 3 and window size of 10 for Oct. 4. 419 This is because a large window size increases the strength of the decoupling and removes the temporal variations of release 420 rates more completely. However, a large window size leads to increased computational cost. Because the errors vary in a 421 limited range, a medium window size provides a better balance between accuracy and computational cost. Feature selection 422 improves the results for medium and small window sizes, but may have less effect with large window sizes. This tendency 423 implies that it is more appropriate to apply feature selection with medium window sizes than with large window sizes, as is 424 done in this study.









427 **3.4.3 Sensitivity to the feature type**

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Figure 11 compares the results obtained with different feature types. For Oct. 3, the localization errors are quite low when using only the time-domain features for the reconstruction; indeed, the errors are only slightly larger than when using all the features. In contrast, the results obtained using only the frequency-domain features exhibit noticeably larger errors, indicating that the time-domain features make a greater contribution to the results for Oct. 3. For Oct. 4, the mean localization errors are similar when using the features of either the time or frequency domain, but the error range is higher when the frequencydomain features are used. In addition, the errors of both single-domain-feature results are higher than those of the all-feature results, indicating that both feature types should be considered to ensure accurate and robust source localization.



435





436 **Figure 11.** Sensitivity to feature type. (a) Oct. 3; (b) Oct. 4.

437 **3.4.4** Sensitivity to the number and combination of observation sites

438 Figure 12 compares the results obtained with different numbers and combinations of observation sites. The results indicate 439 that the localization error may be more sensitive to the position of the observation site than to the number of sites included. 440 For both days, the results obtained using a subset of sites produce the lowest error level, i.e. Site ABD on Oct. 3 and Site BD 441 on Oct. 4. These results are completely consistent with the feature importance results in Fig. 4. The implication is that sampling 442 a plume at multiple locations with correspondingly different activity concentrations is more likely to result in a better 443 reconstruction because the environmental variability is more fully captured through direct observations. For Oct. 3, source 444 localization with Site B always produces lower error levels, though Site B is farthest away in the crosswind direction. However, the results on Oct. 4 do not exhibit the same phenomenon, mainly because almost all observation sites are parallel to the 445 446 downwind direction. In addition, the single-site results can also achieve low error levels, such as Site B on Oct. 3 and Site A 447 on Oct. 4. Feature selection reduces the mean error level in most test cases. These results indicate that the proposed method 448 may achieve satisfactory results with different numbers and combinations of observation sites. It also remains effective with 449 very few observation sites, on the condition that the observations sufficiently capture the plume.







451 **Figure 12.** Sensitivity to the number and combination of observation sites. (a) Oct. 3; (b) Oct. 4.

452 **4. Conclusions**

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453 In this study, we relaxed the unrealistic constant-release assumption of source reconstruction. Instead, we took advantage of 454 the fact that most atmospheric radionuclide releases have a spatially fixed source and thus the release rate mainly influences 455 the peak values in the temporal observations. Based on this, a more general spatiotemporally decoupled source reconstruction 456 method was developed to estimate non-constant releases. Decoupling was achieved by applying a temporal sliding-window 457 average filter to the observations. This filter reduces the influence of temporal variations in release rates on the observations, 458 so that the relative spatiotemporal distribution of the filtered observations is dominated by the source location and known 459 meteorology. A response feature vector was extracted to quantify the long-term temporal response trends at each observation 460 site, involving tailored indicators of both the time and frequency domains. The XGBoost algorithm was used to train a machine 461 learning model that links the source location to the feature vector, enabling independent source localization without knowing 462 the release rates. With the retrieved source location, the detailed temporal variations of the release rate were determined using the PAMILT algorithm. Validation was performed against the two-day SCK-CEN ⁴¹Ar field experimental data, and the results 463 464 demonstrate that the proposed method successfully removes the influence of temporal variations in release rates across observations and accurately localizes the source. Source localization was achieved with deviations of only 4.68% and 4.71% 465





466 on Oct. 3 and Oct. 4, respectively, representing reductions of 9.42% and 9.59%, respectively, compared with the results from 467 a recent correlation-based method and 7.20% and 8.12%, respectively, compared with the results from the Bayesian method. In terms of the release rate, the PAMILT method reconstructed the temporal variations, peak, and total release with high 468 469 accuracy, thus avoiding the unrealistic oscillations given by the Bayesian estimate. The proposed method also exhibited smaller 470 uncertainty ranges in terms of the source location and total release than the Bayesian method, and avoided the multimodal 471 distribution of the Bayesian method. Sensitivity analyses revealed that the proposed method exhibits stable decoupling and 472 localization performance with different parameters and remains effective with only a single observation site, as long as the 473 selected site is at an appropriate position. These results demonstrate that spatiotemporally decoupled source reconstruction is 474 feasible and achieves satisfactory accuracy in the non-constant-release scenario, thereby providing a promising framework for 475 reconstructing atmospheric radionuclide releases.

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477 Code and data availability. The code and data for the proposed method can be downloaded from Zenodo 478 (https://doi.org/10.5281/zenodo.10200141). More recent versions of the code and data will be published on GitHub.com 479 (https://github.com/rocket1ab/Source-reconstruction-of-non-constant-atmospheric-radionuclide-releases, last access: 23 480 November 2023). The implementation is provided in Python, and the instruction file is also available in the provided link.

481

Author contributions. YX conducted the source reconstruction tests and wrote the manuscript draft; SF provided guidance on
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485 *Competing interests.* The authors have declared that they have no conflict of interest.

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