

# 1 **P3D-BRNS v1.0.0: A Three-dimensional, Multiphase,** 2 **Multicomponent, Pore-scale Reactive Transport Modelling** 3 **Package for Simulating Biogeochemical Processes in Subsurface** 4 **Environments**

5  
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## 13 **Abstract**

14 The porous microenvironment of soil offers various environmental functions which are governed by  
15 physical and reactive processes. Understanding reactive transport processes in porous media is essential  
16 for many natural systems (soils, aquifers, aquatic sediments or subsurface reservoirs) or technological  
17 processes (water treatment, or ceramic and fuel cell technologies). In particular, in the vadose zone of the  
18 terrestrial subsurface the spatially and temporally varying saturation of the aqueous and the gas phase  
19 leads to systems that involve complex flow and transport processes as well as reactive transformations of  
20 chemical compounds in the porous material. To describe these interacting processes and their dynamics at  
21 the pore scale requires a well-suited modelling framework accounting for the proper description of all  
22 relevant processes at a high spatial resolution. Here we present P3D-BRNS as a new open-source  
23 modelling toolbox harnessing the core libraries of OpenFOAM and coupled externally to the  
24 Biogeochemical Reaction Network Simulator (BRNS). The native OpenFOAM Volume of Fluid solver is  
25 extended to have an improved representation of the fluid-fluid interface. The solvers are further  
26 developed to couple the reaction module which can be tailored for a specific reactive transport simulation.  
27 P3D-RBNS is benchmarked against three different flow and reactive transport processes; 1) fluid-fluid  
28 configuration in a capillary corner, 2) mass transfer across the fluid-fluid interface and 3) microbial  
29 growth with a high degree of accuracy. Our model allows for simulation of the spatio-temporal  
30 distribution of all bio-chemical species in the porous structure (obtained from  $\mu$ -CT images), for  
31 conditions that are commonly found in the laboratory and environmental systems. With our coupled  
32 computational model, we provide a reliable and efficient tool for simulating multiphase, reactive transport  
33 in porous media.

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## 34 1 Introduction

35 Subsurface environments (soils, aquifers, aqueous sediments) are (typically) porous media host a  
36 multitude of biogeochemical processes and interactions and provide different versatile ecosystem  
37 functions (e.g. C sequestration, compound degradation, nutrient retention, provision of food, fibers and  
38 fuel, habitat for organisms, water retention and purification, etc. (Baveye et al., 2016 ). These processes  
39 are controlled by various biological (e.g., microbial abundance and activity), chemical (e.g., distribution  
40 of dissolved and volatile species, mineral composition and surface properties of the solid matrix) and  
41 physical (e.g., porous structure and permeability, water saturation) properties of the system. These  
42 features create a complex web of interactions, the magnitude and effectiveness of which change  
43 dynamically in space and time (Graham et al., 2014). Microbial communities, for example, and their  
44 metabolic capacity are considered to be directly related to energy and matter fluxes (Thullner et al., 2007)  
45 which are in turn, governed by pore arrangements and their connectivity. Along with other environmental  
46 factors this can also modify various properties of the porous media (e.g. by biomass accumulation on pore  
47 walls (Thullner, 2010), or mineral dissolution or precipitation (Meakin and Tartakovsky, 2009), which in  
48 turn are altering the conditions for biogeochemical processes, too.

49 In soils (or more generally the vadose zone) the dynamically varying distribution of the aqueous and  
50 gaseous phase leads to specifically complex and variable constraints for biogeochemical processes. In the  
51 past, obtaining (bio)chemical and microbiological information at the pore level was neither economically  
52 nor logistically a feasible option (Baveye et al., 2014). Also for the sake of applicability, traditionally,  
53 researchers had more tendency to look for ~~macro-scale~~-~~Darcy-scale~~ solutions to tackle environmental  
54 issues (White and Brantley, 2003). The ~~Darcy-scale~~ ~~macro-scale~~-view (experimental, theoretical or a  
55 mixture of both) serves well the purpose of practical applicability (White and Brantley, 2003), but for  
56 example, in the context of microbially mediated degradation processes in the vadose zone, it fails to  
57 provide insights on the driving mechanisms, as it overlooks important contributing factors, such as the  
58 tortuous porous structure/pathways open to the transport of (bio-)geochemical species, non-uniform  
59 distribution of water and air phases, as well as the nonlinear ~~relation~~-dependency between changes of the  
60 local nutrient and biomass concentrations and the bulk concentrations of (bio)geochemical species.~~change~~  
61 of in local biomass concentration and the its relation with bulk nutrient concentration. Evidence at  
62 microscopic level has shown that biological activity and evolution are more locally organized (Kuzyakov  
63 and Blagodatskaya, 2015) where ~~Darcy-scale~~ ~~macroscopic~~-studies lead to loss of crucial information.  
64 This has motivated the development of sophisticated physics-based models implementing all aspects of  
65 hydrological, geochemical and biological processes involved in microbial growth and evolution.

66 Reactive Transport Models (RTMs) are a class of mathematical models that have been applied  
67 extensively to study biogeochemical systems for about four decades (Parkhurst and Appelo,  
68 1999;Thullner et al., 2005;Thullner and Regnier, 2019;Meile and Scheibe, 2019). There is a long list of  
69 principal factors and mechanisms governing biogeochemical reactions at the pore scale. Numerically,  
70 these processes can be defined and solved either by fully (global) implicit approaches or by separating  
71 and solving different components once at a time. For the continuum scale a wide range of reactive  
72 transport models exist which allow for the simulation of biogeochemical processes (Steeffel et al., 2015b).  
73 In turn, at the pore scale, models combining the simulation of flow, transport and reactive  
74 (biogeochemical) processes are scarce, and existing model developments are often driven by specific  
75 research questions and/or are subject to severe simplifications in the description of the pore space  
76 (Golparvar et al., 2021). Integrated models- explicitly capturing simultaneously the structural properties of  
77 the soil at the microscale-pore -scale, the resulting multiphase flow and multispecies reactive transport are  
78 hardly available (Tian and Wang, 2019).

79 Recently, new frontiers of pore-scale RTMs are emerging with the advances in computational power as  
80 well as with huge improvements in imaging techniques. The latter includes (e.g., ~~in-the~~ static and  
81 dynamic scanning of porous structure as well of as fluids' distribution (Schlüter et al., 2019) or ~~in-the~~  
82 detectiong of bacterial distributions in soil using ~~e.g.,~~ catalyzed reporter deposition with fluorescence in  
83 situ hybridization (CARD-FISH) technique (Schmidt et al., 2015). Direct Numerical Models (DNMs) are  
84 becoming the nexus of next generation of RTMs as they represent the porous structure in a fully explicit  
85 manner (directly obtained from soil samples, digitized and fed into RTMs) in addition to offering a more  
86 flexible coupling of different components of reactive transport models (Baveye et al., 2018;Raeini et al.,  
87 2012;Li et al., 2010;Yan et al., 2016). Another advantage of using DNMs is that they offer a great deal of  
88 flexibility in considering settings and conditions that are experimentally impossible to impose (Tian and  
89 Wang, 2019).

90 In this work, we introduce the pore-scale RTM package P3D-BRNS explicitly involving the structure and  
91 topology of the pore space, the co-existence/co-flow of both the aqueous and the gaseous phase, the  
92 advective-diffusive transport of species in each phase, and an arbitrary set of reactive processes controlled  
93 by kinetic rate laws or thermodynamic constraints. The fluid flow field is updated via solving the Navier-  
94 Stokes (NS) equations (Patankar and Spalding, 1972). The Volume-Of-Fluid (VOF) approach is adopted  
95 to account for different phase distribution (Hirt and Nichols, 1981). The transport of chemical species is  
96 considered via solving the Advection-Diffusion-Reaction Equation (ADRE), where the concentration  
97 jump for soluble/volatile compounds across the fluid/fluid interface is modelled via the Continuous  
98 Species Transfer (CST) method (Haroun et al., 2010). Reactive processes are defined and simulated

99 externally via coupling the flow and transport model to the BRNS (Biogeochemical Reactions Network  
 100 Solver) package (Regnier et al., 2002; Aguilera et al., 2005). The model structure is introduced and the  
 101 model performance is shown and compared with analytical counterparts. The model capabilities are  
 102 depicted for a fully three-dimensional case.

## 103 2 Mathematical Formulation

104 The entire numerical domain ( $\Omega$ ) can be decomposed to two main sub-regions: solid space ( $\Omega_S$ ) and void  
 105 space ( $\Omega_\theta$ ). The void space is further divided into aqueous phase ( $\Omega_{\theta,aq}$ ) and gaseous phase ( $\Omega_{\theta,gs}$ )  
 106 which are partitioned by the fluid-fluid interface ( $\Omega_I$ ). The overall domain is bounded externally between  
 107 inlet ( $\Omega_{in}$ ) and outlet ( $\Omega_{out}$ ) boundaries, which allow for inflow/outflow of different phases and chemical  
 108 species, as well as no-flux boundaries resembling physical walls, where nothing is allowed to leave or  
 109 enter the domain ( $\Omega_{wall}$ ). The domain is limited internally by no-flow boundaries where solid space and  
 110 void space intersect (i.e. solid surface,  $\Omega_\theta \cap \Omega_S = \Omega_{wall}$ ).

### 111 2.1 Fluid Flow: Governing Equations

112 Evolution of a single/multi-phase, isothermal, incompressible, immiscible fluid(s) can be expressed by  
 113 basic conservation principles. These can be formulated into a single-field formalism (Hirt and Nichols,  
 114 1981):

$$\nabla \cdot \mathbf{u} = 0 \quad \text{in } \Omega_\theta \quad (1)$$

$$\rho \left( \frac{\partial \mathbf{u}}{\partial t} + \mathbf{u} \cdot \nabla \mathbf{u} \right) = -\nabla P + \nabla \cdot \boldsymbol{\tau} + \rho \mathbf{g} + \mathbf{F}_\sigma \quad \text{in } \Omega_\theta \quad (2)$$

115 where  $\mathbf{u}$  is the vector of velocity field,  $\rho$  is the fluid density,  $P$  is the pressure,  $\mathbf{g}$  is the gravitational  
 116 vector, and  $\boldsymbol{\tau}$  is the stress tensor which can be defined as  $\boldsymbol{\tau} = 2\mu \mathbf{S} = 2\mu(0.5[(\nabla \mathbf{u}) + (\nabla \mathbf{u})^T])$  with  $\mu$  as  
 117 the fluid viscosity.  $\mathbf{F}_\sigma$  denotes the interfacial tension force, which is nonzero only when two or more  
 118 phases (excluding solid) are available. It is safe in using the incompressible form of the Navier-Stokes  
 119 equation for low Mach and Reynolds numbers

120 In case of simultaneous flow of two different phases (e.g. flow of air and water under unsaturated conditions  
 121 –Figure 1: Illustration of a porous medium at the pore scale with one fluid invading the other (on the left).  
 122 Mathematical representation of the phase saturation in the computational cells around the interface (on  
 123 the right). The dashed line shows the actual location of the interface while values in each cells show the  
 124 amount of water saturation relevant to topology of the dashed line. Black dots represent volatile  
 125 compounds able to cross the fluid-fluid interface; green dots represent non-volatile compounds restricted

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126 to the transport in the aqueous phase. Figure 1), their locations and distribution are represented via  
 127 introducing an indicator function,  $\alpha$ , taking values within the range  $[0, 1]$ . The first continuous fluid is  
 128 marked as  $\alpha = 1$ , the second fluid is denoted as  $\beta = 1 - \alpha$  and for the transition from one fluid to the other  
 129 (i.e. the interface,  $\Omega_I$ ),  $\alpha$  varies between 0 and 1. Fluid density and viscosity in each grid cell is then  
 130 calculated from a linear interpolation of this indicator function:

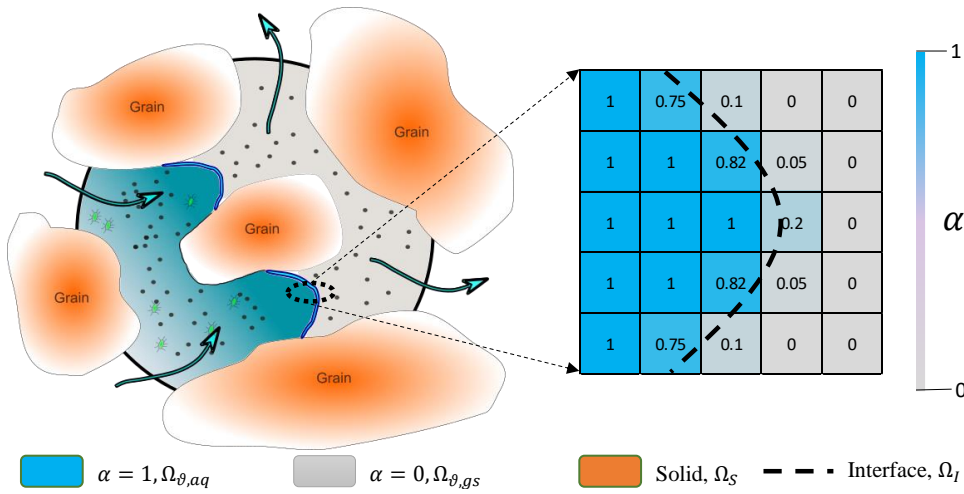
$$\rho = \alpha\rho_1 + (1 - \alpha)\rho_2 \quad \text{in } \Omega_\theta \quad (3)$$

$$\mu = \alpha\mu_1 + (1 - \alpha)\mu_2 \quad \text{in } \Omega_\theta \quad (4)$$

131 A mass conservative boundary condition at the fluid-fluid interface is written as:

$$[[\rho_i(\mathbf{u}_i - \mathbf{w}) \cdot \mathbf{n}_{\Omega_I}]] = 0 \quad \text{in } \Omega_I \quad (5)$$

132



**Figure 1:** Illustration of a porous medium at the pore-scale microscale with one fluid invading the other (on the left). Mathematical representation of the phase saturation in the computational cells around the interface (on the right). The dashed line shows the actual location of the interface while values in each cells show the amount of water saturation relevant to topology of the dashed line. Black dots represent volatile compounds able to cross the fluid-fluid interface; green dots represent non-volatile compounds restricted to the transport in the aqueous phase.

133 with  $\rho_i$  as the density of  $i^{\text{th}}$  fluid,  $\mathbf{u}_i$  as the velocity of  $i^{\text{th}}$  fluid,  $\mathbf{w}$  as the velocity of the interface,  $\mathbf{n}_{\Omega_I}$  as

134 the normal vector to the interface ( $\Omega_I$ ) pointing from the invading phase to the displaced one and the

135 brackets showing the jump condition at the interface. Individual velocities,  $\mathbf{u}_i$ , and the interface velocity

136  $\mathbf{w}$ , are not directly calculated but furthermore averaged to derive the global mass conservation equation

137 that is used for numerical discretization (for full derivation, consult with Graveleau et al., 2017).

138 In the context of the Finite Volume Method (FVM), discretization of the physical domain produces a  
 139 finite subset of discrete volumes (taking the shape of a polyhedral). The key implication of the Volume of  
 140 Fluid method is to define and solve for global variables, rather than having one equation for each variable  
 141 in each phase. Hence, the idea is to transform the integro-differential equations into their global versions  
 142 by averaging them over each cell volume (Whitaker, 2013). For multiphase systems, after a few steps of  
 143 linearization and approximation (see Hirt and Nichols (1981) for a detailed derivation), the Volume of  
 144 Fluid formulation of the momentum equation ~~(2)(2)~~ is obtained as:

$$\rho \left( \frac{\partial \bar{\mathbf{u}}}{\partial t} + \bar{\mathbf{u}} \cdot \nabla \bar{\mathbf{u}} \right) = -\nabla P + \nabla \cdot \mu (\nabla \bar{\mathbf{u}} + (\nabla \bar{\mathbf{u}})^t) + \rho g + F_\sigma \quad \text{in } \Omega_\theta \quad (66)$$

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145 with having  $\bar{\mathbf{u}}$  as the global averaged velocity vector. For the sake of simplicity, we drop the “average”  
 146 notation from the global velocity vector (i.e. will refer to  $\bar{\mathbf{u}}$  as  $\mathbf{u}$ ), for the rest of this paper. The pressure  
 147 gradients (and Reynolds numbers) considered in our simulations are in the range that render changes in  
 148 the gas compressibility negligible.  
 149 Since we are dealing with only two fluids, index  $i$  in  $\mathbf{u}_i$  takes only two values:  $\alpha$  and  $\beta$  - one for each  
 150 phase. A global, mass conservative, advection equation is used to describe the evolution of the indicator  
 151 function:

$$\frac{\partial \alpha}{\partial t} + \nabla \cdot (\alpha \mathbf{u}) + \nabla \cdot (\alpha(1 - \alpha) \mathbf{u}_c) = 0 \quad \text{in } \Omega_\theta \quad (7)$$

152 where  $\alpha$  indicates the volume fraction of phase 1,  $\mathbf{u}_c = \mathbf{u}_\alpha - \mathbf{u}_\beta$  is the vector of the compressive velocity  
 153 with  $\mathbf{u}_{\alpha,\beta}$  as velocity vector of phase  $\alpha$  and  $\beta$  right on the edge of the interface (detailed explanation on  
 154 deriving eq. ~~(7)(7)~~ can be found in the Supplementary Information Error! Reference source not found.).  
 155 The latter term in equation (7) is active only in the presence of an interface. It is derived from mass  
 156 conservation equation written for phase  $\alpha$ , which computationally helps with maintaining  
 157 stiffness/sharpness of the interface. Sharpening interface means having the interface span over a fewer  
 158 number of computational grids.  $\mathbf{u}_c$  is the vector of compressive velocity on each face of all computational  
 159 grids. Since we don't solve for the velocity field of each phase individually, a direct calculation of  $\mathbf{u}_c$  is  
 160 not possible. However, we can rather take an indirect approach for computing  $\mathbf{u}_c$  as follows:

$$\mathbf{u}_c = \min(c_\alpha |\mathbf{u}|, \max(|\mathbf{u}|)) \frac{\nabla \alpha}{|\nabla \alpha|} \quad \text{in } \Omega_l \quad (8)$$

161 In equation ~~(8)(8)~~,  $c_\alpha$  is a compression coefficient providing some level of control over how wide the  
 162 interface spans. The max function operates on the magnitude of unit velocity vector calculated on all the  
 163 faces of a computational grid. To counteract the numerical diffusion and avoid the spread of interface  
 164 over several computational grids, values of  $c_\alpha > 1$  provide an enhanced/sharper interface. A where a

165 value of  $c_\alpha = 0$  gives no compression of the interface, ~~whereas values of  $c_\alpha > 1$  provide an~~  
 166 ~~enhanced/sharper interface~~ (Graveleau et al., 2017). In simulation scenarios introduced in this paper,  $c_\alpha$   
 167 has been assigned the value of 1, ~~unless stated otherwise~~~~unless otherwise is stated~~.

168 To calculate the interfacial tension force,  $\mathbf{F}_\sigma$ , (Brackbill et al., 1992) have introduced a Continuum  
 169 Surface Force (CSF) which requires computing the curvature of the interface:

$$\kappa = \nabla \cdot n_{\Omega_I} \quad \text{in } \Omega_I \quad (99)$$

170 with having  $\kappa$  as mean interface curvature in each computational grid, and  $n_{\Omega_I}$  as the interface unit  
 171 normal vector defined as:

$$n_{\Omega_I} = \frac{\nabla \alpha}{|\nabla \alpha|} \quad \text{in } \Omega_I \quad (104)$$

172 Given the curvature, the interfacial tension force can be ~~computed~~approximated as:

$$\mathbf{F}_\sigma = \sigma \kappa \nabla \alpha \quad \text{in } \Omega_I \quad (111)$$

173 where  $\sigma$  is the surface tension between two fluids (derivation of this approximation can be found in  
 174 Brackbill et al., 1992).

## 175 2.2 Reactive Transport: Governing Equations

176 Concentrations of mobile species are affected by advection (i.e. transport with the moving fluid),  
 177 molecular diffusion and reactive transformation. Also, in case of having two fluids simultaneously in the  
 178 system, different species can cross the fluid-fluid interface, causing local fluctuations in concentration  
 179 values. In general to account for all the changes in species concentrations, the ADRE for biogeochemical  
 180 reactive components can be written as:

$$\frac{\partial C_i}{\partial t} + \nabla \cdot (C_i \mathbf{u}) = -\nabla \cdot (J_{d,i} + J_{m,i}) + R_i \quad \text{in } \Omega_g \quad (121)$$

181 where  $J_{d,i}$  is the molecular diffusive flux of component  $i$ ,  $J_{m,i}$  is the mass flux of component  $i$  due to mass  
 182 transfer across the fluid-fluid interface and  $R_i$  accounts for the changes in concentration of component  $i$   
 183 due to reactions. Molecular diffusion follows Fick's law:

$$J_{d,i} = -D_i \nabla (C_i) \quad \text{in } \Omega_g \quad (131)$$

184 where  $D_i$  is the diffusion coefficient of species  $i$ . At the interface, the assumption of thermodynamic  
 185 equilibrium implies equality of chemical potentials. Given the condition that liquid concentration of  
 186 component  $i$  is proportional to the partial pressure of the species in the secondary phase (e.g. gas, oil or  
 187 minerals), a partitioning relationship such as Raoult or Henry's law (Danckwerts and Lannus, 1970) can  
 188 be established to relate species concentrations on both sides of the interface:

$$C_{i,\beta} = H_i C_{i,\alpha} \quad \text{in } \Omega_l \quad (14) (14)$$

189 with  $C_{i,\alpha}$  as concentration of species  $i$  in phase  $\alpha$ ,  $H_i$  as Henry's constant of species  $i$  and  $C_{i,\beta}$  as  
 190 concentration of species  $i$  in phase  $\beta$ . Depending on if a given compound's concentration in the aqueous  
 191 phase or the gaseous phase is multiplied by the Henry's coefficient, ~~(14)~~(14), the definition of Henry's  
 192 constant switches between the solubility or volatility for that compound (i.e.  $H_i^{solubility} = 1/H_i^{volatility}$ )  
 193 (Sander, 2015). Unless otherwise stated, the volatility concept of Henry's law is adopted in order to  
 194 define the concentration relationship of a given compound across the fluid-fluid interface ~~(14)~~(14). The  
 195 concentration field around the fluid-fluid interface (where  $\nabla\alpha \neq 0$ ) at equilibrium, for any values of  $H \neq$   
 196 1, is discontinuous which imposes the additional flux,  $J_{m,i}$ , to satisfy the concentration jump across the  
 197 interface. Hence the mass transfer flux,  $J_{m,i}$ , can be derived within the VOF framework (i.e. CST) as  
 198 follows (Haroun et al., 2010):

$$J_{m,i} = -D_i \frac{1 - H_i}{\alpha + (1 - \alpha)H_i} C_i \nabla\alpha \quad \text{in } \Omega_\theta \quad (15) (15)$$

199 It is noteworthy that few assumptions and volume averaging methods are implemented to derive equation  
 200 ~~(15)~~(15), which readers are encouraged to check the references for more details. The diffusion coefficient  
 201 is calculated from harmonic interpolation:

$$D_i = \frac{1}{\frac{\alpha}{D_{i,A}} + \frac{1 - \alpha}{D_{i,B}}} \quad \text{in } \Omega_\theta \quad (16) (16)$$

202 where  $D_{i,\alpha-\beta}$  is the diffusion coefficient of species  $i$  in phase  $\alpha$  and  $\beta$  respectively.

203 Simulated reactions include kinetically as well as thermodynamically constrained reactions. For a  
 204 kinetically constrained reaction  $j$  the reaction rate  $r_j = f(C_1, \dots, C_n)$  is needed while for a  
 205 thermodynamically constrained reaction  $k$  the equilibrium conditions defined by a law of mass action  
 206  $M_k = f(C_1, \dots, C_n)$  is needed with  $M_k$  as equilibrium constant. These equations can be of arbitrary form and  
 207 the resulting reaction network defines the term  $R_i$  in equation ~~(12)~~(12) (Aguilera et al., 2005; Regnier et al.,  
 208 2002). For immobile species concentration changes are only due to reactive processes.

### 209 2.3 Boundary Conditions (BCs)

210 There are various types of boundary conditions, corresponding the real physical conditions, most of which  
 211 can be derived from two basic types:

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- 212 • Dirichlet boundary (fixed value) which relates the value of a variable at a given geometric  
213 location to a constant value; e.g.  $C_i = 1M$  in  $\Omega_{in}$  meaning a constant 1 molar concentration of  
214 component  $i$  at the boundary,
- 215 • von Neumann boundary (fixed gradient) which provides the value of a variable's gradient at the  
216 face of the boundary cell; e.g.  $\partial_n p = 0$  in  $\Omega_{wall}$  giving a zero pressure gradient on the wall.

217 In general, our model can apply any of these basic boundary conditions to any scalar or vector variables  
218 such as pressure, velocity/flux, concentration of volume of fluid fields, but one needs to assure that the  
219 imposed BC(s) are both compatible and they reflect the correct physical boundary conditions. For  
220 example, for velocity/flux-pressure coupling, a Dirichlet (i.e. constant) boundary for flux at the inlet can  
221 be coupled with either 1) fixed discharge velocity/flux and zero gradient (i.e. von Neumann) pressure at  
222 both inlet and outlet, or 2) a constant pressure head at the inlet and atmospheric pressure at the outlet with  
223 zero gradient velocity/flux at both ends, or 3) fixed values of pressure and velocity/flux at one end and  
224 zero gradient at the other end. In the beginning of section 22, typical composition and configuration of an  
225 arbitrary computational domain is described. Inlet, outlet and impermeable boundaries are amongst the  
226 most common types that one might face. Inlet BC means for the direction of fluid flux to be pointing  
227 inwards (i.e. into the domain) while for the outlet, the direction of the flux should be outwards. Also for  
228 the impermeable wall, zero-orthogonal fluxes need to be satisfied. Either Dirichlet, von Neumann or a  
229 mixture of both can be used at a particular boundary. Mathematical translation and implementation of  
230 these boundaries are provided in the next section. Time-dependent BCs (e.g. cyclic or seasonal  
231 water/species influx) are also readily available to be applied, but never been used in this work. Unless  
232 otherwise stated, boundary conditions that have been imposed on each section of the computational  
233 domain are described as follows:

234 **At impermeable boundaries ( $\Omega_{wall}$ ):** Physical wall implies no flux perpendicular to the normal vector  
235 to its surface. No slip BC is an appropriate BC for the velocity field on the wall. In general, they all can  
236 be written as:

$$\partial_n C_i = 0, \quad \mathbf{u}_{x,y,z} = 0, \quad \partial_n p = 0, \quad \partial_n \alpha = 0 \quad \text{in } \Omega_{wall} \quad (17+7)$$

237 For the velocity field, on the wall, a slip boundary condition is also available to be applied.

238 **At inlet/outlet boundaries ( $\Omega_{in}, \Omega_{out}$ ):** Concentration of reactants, products and inert tracers are set to  
239 fixed values at inlet, while they are allowed to leave the domain at outlet with zero gradient boundary  
240 condition. Constant flowrate with zero pressure gradient is applied at inlet and an atmospheric pressure  
241 (fixed value) with zero velocity gradient is set at outlet. Also in case of two-phase flow, the invading

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242 phase is set to enter from inlet at fixed value and exits from outlet with zero gradient BC. Mathematically,  
 243 they can be expressed as:

$$C_i \geq 0, \quad \mathbf{u} = \text{const.}, \quad \partial_n \mathbf{p} = 0, \quad \alpha = \text{const.} \quad \text{in } \Omega_{in} \quad (18)_{18}$$

244 together with:

$$\partial_n C_i = 0, \quad \partial_n \mathbf{u}_n = 0, \quad \mathbf{p} = 0, \quad \partial_n \alpha = 0 \quad \text{in } \Omega_{out} \quad (19)_{19}$$

245 with  $u_n$  as the normal velocity vector. While we have mostly applied equations (18)(18) and (19)(19) for  
 246 designing an inlet/outlet duo, other formats, such as defining a pressure head (plus zero gradient velocity)  
 247 on the inlet in combination with either constant exit pressure or constant discharge rate, are readily  
 248 available to implement as well.

249 **At the fluid-fluid-solid contact line ( $\Omega_{i3}$ ):** At the fluid-fluid-solid contact line, in case of no interactions  
 250 or no reaction of any chemical species with the solid, the boundary condition at the triple point is derived  
 251 to be:

$$\nabla C_i \cdot \mathbf{n}_s = \frac{H_i - 1}{\alpha H_i + (1 - \alpha)} C_i \nabla \alpha \cdot \mathbf{n}_s \quad \text{in } \Omega_l \quad (20)_{20}$$

252 with  $\mathbf{n}_s$  as the normal vector to the solid surface (Graveleau et al., 2017). Also, the concept of contact  
 253 angle is applied by making the following modification to the interface normal vector:

$$\mathbf{n}_{\Omega_{i3}} = \cos\theta \mathbf{n}_s + \sin\theta \mathbf{t}_s \quad \text{in } \Omega_{i3} \quad (21)_{21}$$

254 where  $\mathbf{n}_s$  is normal vector and  $\mathbf{t}_s$  is the tangential vector to the solid surface (Brackbill et al., 1992). At  
 255 the triple point, i.e. fluid-fluid-solid interface,  $\mathbf{n}_{\Omega_{i3}}$  is used for normal vector to the interface. CSF,  
 256 though, has been reportedly generating non-physical spurious currents (Scardovelli and Zaleski, 1999).  
 257 For this, many have tried to eliminate/mitigate this issue by explicit representation of the interface either  
 258 via using the Geometric VOF method (Popinet, 2009) or coupled Level-set (LS) VOF functions  
 259 (Albadawi et al., 2013). Geometric VOF is quite suitable for structured grids, but for porous structures  
 260 with highly unstructured grids, the calculations can become quite complicated. Alternatively, Raeini et al.  
 261 (2012) suggested filtering the capillary forces parallel to the interface, which can significantly reduce the  
 262 non-physical velocities. In short, the modifications they proposed and which are used here are: 1)  
 263 smoothing the indicator function to have a better measure of the interface curvature, 2) sharpening the  
 264 indicator function for computation of the interfacial tension force, 3) filtering the capillary pressure force  
 265 parallel to the interface, and 4) filtering capillary fluxes based on the capillary pressure gradient (for full  
 266 description of each point, please consult with (Raeini et al., (2012)).

267 To sum up what has been presented so far, we integrated ~~the a)~~ the original *interFoam* solver from the  
 268 OpenFOAM library that only solves for the advection-diffusion transport of two phase flow, with ~~the b)~~

269 [the improved-interface-resolver library from Raeini et al. \(2012\)](#), and ~~added-c~~ [added a scalar transport](#)  
 270 [solver on top of them. Finally, the full-scale advection-diffusion-reaction model of the biogeochemical](#)  
 271 [species is attained by coupling this to an external reaction-network solver, which is explained in the](#)  
 272 [section 2.4 below.](#)

## 273 2.4 Numerical Formulation

274 The mass conservation (eq. [\(14\)](#)), momentum (NS - eq. [\(22\)](#)) and indicator function (eq. [\(77\)](#)) equations are  
 275 all implemented within the open source computational fluid dynamics (CFD) package, OpenFOAM  
 276 (Greenshields, 2015). OpenFOAM utilizes the finite volume methodology (FVM), a common choice for  
 277 CFD problems as FVM works only with conservative flux evaluation at each computational cell's  
 278 boundaries, making it robust in handling nonlinear transport problems. Also all the differential equations  
 279 mentioned before are first written in their integral form over each cell volume and then converted to the  
 280 surface summations using Green's Theorem.

281 ~~In this work,~~ [The original two-phase \(VOF\) flow solver, i.e. \*interFoam\*, is modified to construct our](#)  
 282 [biogeochemical reactive transport package. The momentum equation \[\\(22\\)\]\(#\) is linearized in a semi-discrete](#)  
 283 [form as:](#)

$$A_d \mathbf{u} = \mathbf{H}(\mathbf{u}) - \nabla P + \mathbf{F}, \quad (222)$$

284 where  $A_d$  holds the diagonal elements of the coefficient matrix,  $\mathbf{H}(\mathbf{u})$  contains off-diagonal elements of  
 285 the coefficient matrix including all source terms, and  $\mathbf{F}$  entails any body forces (interfacial tension force  
 286 only in this case). Temporal discretization is handled via the first order Euler method while spatial  
 287 discretization is managed via second order finite volume schemes. Convection terms of the momentum  
 288 equation and indicator function [\(7\)](#)~~(eq. 7)~~ are computed using a bounded self-filtered central differencing  
 289 (SFCD) scheme (based on Gauss's theorem). Rearranging equation [\(22\)](#)~~(22)~~ for velocity and imposing the  
 290 continuity equation [\(14\)](#), the following linear pressure equation can be obtained:

$$\sum_f \frac{S_f}{\langle A_d \rangle_f} \nabla_f^\perp P = \sum_f \left( \left\langle \frac{\mathbf{H}(\mathbf{u})}{A_d} \right\rangle_f \cdot |S_f| + \frac{\varphi_{F,f}}{\langle A_d \rangle_f} \right). \quad (232)$$

291  $S_f$  in equation [\(23\)](#)~~(23)~~ denotes the outward area-vector of face  $f$ , the notation  $\nabla_f^\perp$  shows face normal  
 292 gradients calculated right on the face centers,  $\langle \cdot \rangle_f$  shows the interpolated values of a face-centered  
 293 parameter from its cell-centered counterpart, and  $\varphi_{F,f}$  is the interfacial force flux term.

294 The velocity-pressure coupling of equations [\(14\)](#) and [\(22\)](#) are solved using Pressure Implicit with  
 295 Splitting of Operators (PISO) (Issa, 1986). PISO embodies a predictor-corrector strategy to  
 296 simultaneously update pressure and velocity within each time step. The resultant system of equations are

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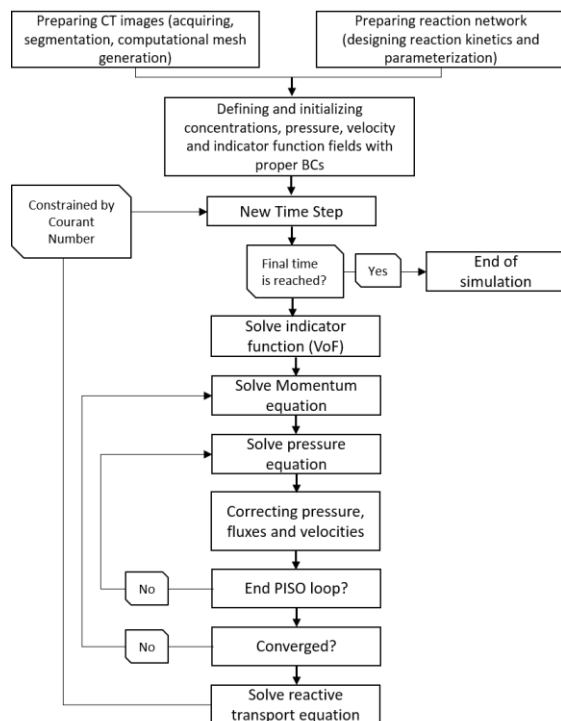
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297 solved on the cell faces and then interpolated back to calculate velocities and pressure at the cell centers.  
298 The coupling of indicator function (eq. (7)) and momentum equation is explicitly defined and solved  
299 right after the PISO step is finished. Within the same time step, transport and reaction of different species  
300 are then solved sequentially – using a Sequential Non-Iterative Algorithm (Steeffel et al., 2015a; Steefel  
301 and MacQuarrie, 1996). Time step size is controlled by introducing a Courant number. Time is discretized  
302 using either Euler or Crank-Nicholson methods and spatial discretization is performed using the Van Leer  
303 second order Total Variation Diminishing scheme (TVD) (van Leer, 1979).

304 The reaction network is built separately and externally solved within the BRNS package - which employs  
305 first order Taylor series expansion terms and uses Newton-Raphson method to iteratively solve the  
306 system of linear equations (Regnier et al., 2002). BRNS utilizes MAPLE programming language to  
307 construct the Jacobian matrix (which contains the partial derivatives of unknown parameters, i.e.  
308 concentrations) and other problem-related data such as rate parameters and translating them to a  
309 FORTRAN package. The FORTRAN code is then compiled to generate shared object (\*.so file) that can  
310 be dynamically called later from the transport solver (Centler et al., 2010). The significance of having  
311 dynamically shared object file is more apparent when running computationally-demanding  
312 cases/scenarios while decomposing and running the application in parallel. BRNS is invoked once the  
313 new concentrations are computed from the transport solver. The updated concentrations from the BRNS  
314 library (i.e. updating concentrations from redox reactions) are then fed back into the transport solver  
315 before moving to the next time step. This process repeats until the final time is reached. This coupling  
316 scheme has been successfully used for other RTM approaches before (Centler et al., 2010; Gharasoo et al.,  
317 2012; Nick et al., 2013). As the reactions are localized, the reaction solver is modular, and OpenFOAM  
318 inherently provides parallelized simulations (via domain decomposition), the P3D-BRNS can easily be  
319 used to model larger systems. To achieve higher performance, it is recommended to utilize physical cores  
320 than using hyper-threading. The parallelization of our model strengthen its scalability in the sense of the  
321 size (pore scale or Darcy scale) of the simulated system. However, in terms of upscaling (e.g., from the  
322 pore scale to the Darcy scale) an intermediate step would be required depending on the complexity of the  
323 processes that are involved and on the size of the domain.

324 Prior to run simulations, the physical settings of the domain are required to be specified; i.e. the physical  
325 geometry of the pore space with proper boundaries and the meshing scheme should be designed.  
326 OpenFOAM provides a basic utility for defining boundaries as well as mesh generation which are  
327 translatable by the OpenFOAM engine. Any other meshing software/freeware can be freely used as long  
328 as an OpenFOAM-compatible format for the meshed file can be created. The overall workflow required  
329 to build and run a case/scenario is summarized in [Figure 2](#).

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331 **Figure 2:** Full solution procedure to simulate a reactive transport process at its fullest complexity.

### 332 3 Model Performance

333 The presented reactive transport model is designed 1) to capture real world pore structures in up to three  
 334 dimensions, 2) to explicitly simulate the transient distribution of a gas and a liquid phase within the entire  
 335 pore space and 3) to simulate a full set of advection-diffusion-reaction mechanisms. Capturing the correct  
 336 curvature of the interface depends heavily on the grid resolution. For a fixed velocity magnitude, higher  
 337 grid resolution, enforces shorter time-step size (from Courant number) for the numerical simulations to  
 338 converge. ~~F~~Also to validate different features of the model various simplified scenarios were used which  
 339 allow the use of analytical expressions as reference for the numerical results. We here show three  
 340 representative test scenarios addressing different features of the model (two-phase flow, mass transfer  
 341 across the fluid-fluid interface and reactive transport) individually. Subsequently, the model capabilities  
 342 are depicted in a final biodegradation scenario making use of the various model features simultaneously.

### 343 3.1 Fluid Configurations

344 In order to test our model's performance in simulating two-phase flow, we have zoomed into a two-  
345 dimensional porous structure and isolated only one single corner taking the shape of an equilateral  
346 triangle. A triangulated mesh is adopted that naturally conforms to the overall shape of the domain.  
347 Initially, two immiscible fluids (one wetting, and one non-wetting, e.g. water and air) are placed in such a  
348 way that their interface forms a straight line (Figure 3, a). The side length of the triangle is 1 mm  
349 with a mesh size of 1  $\mu\text{m}$ . Under thermodynamic equilibrium conditions, the force exerted by the pressure  
350 difference between two phases is countered by the interfacial tension force. This, along with the contact  
351 angle of the non-wetting phase at the wall surface in presence of the wetting phase (e.g. water),  
352 determines the topology of the fluid-fluid interface. For a given corner half-angle, the distance that  
353 wetting phase spreads over the solid surface from the corner vertices (the highlighted section with green  
354 color on Figure 3, b),  $b$ , can be calculated as:

$$b = r \frac{\cos(\theta + \beta)}{\sin(\beta)} \quad (2424)$$

355 with  $r$  as the radius of the interface's curvature,  $\theta$  as the contact angle and  $\beta$  as the corner half-angle  
356 (Blunt, 2017). In order to reach thermodynamic equilibrium, we performed transient, two-phase flow  
357 simulations to compute velocity, pressure and indicator function fields until the triple contact line ( $\Omega_{t3}$ ) is  
358 static. For this, we first divided the equilateral triangle in half, as the problem is symmetrical along the  
359 height of the triangle. The symmetrical plane implies that there is no gradient (of any scalar or vector  
360 field) perpendicular to its surface while the tangential components (of all fields) remain the same. To find  
361 the fluid configuration at equilibrium, we simulated the two-phase flow scenario in two steps. First, we  
362 applied a closed boundary condition on the bottom domain by setting  $\mathbf{u} = 0$  together with  $\partial_n p = 0$ . Also  
363 a closed boundary is imposed on the topmost part of the domain which follows the same BC as the  
364 bottom. This way, the interface is able to reconfigure and reorient itself in order to recreate the imposed  
365 contact angle with the wall, and at the same time, pressure is allowed to build up in both phases to support  
366 the shape of the interface. Then, in order to obtain an equilibrium curvature for the interface, bottom and  
367 top domains are opened. This is achieved by setting the 1) pressure in  $\Omega_{in}$  to the average pressure within  
368 the non-wetting phase, 2) pressure in  $\Omega_{out}$  to the average pressure within wetting phase together with 3)  
369  $\partial_n \mathbf{u} = 0$  on both  $\Omega_{in}$  and  $\Omega_{out}$ . At this stage, we applied a special BC for the indicator function to allow  
370 the fluids to enter or leave the domain at both ends, so that the interface can freely transition to its static  
371 shape. At the inlet ( $\Omega_{in}$ ), the BC for  $\alpha$  is set to switch between  $\partial_n \alpha = 0$ , if the fluid flux is pointing  
372 outwards, and  $\alpha = 0$  if the fluid flux is directed into the domain. Also at the outlet ( $\Omega_{out}$ ), the BC for  $\alpha$   
373 switches between  $\alpha = 1$ , if the fluid flux is inwards, and  $\partial_n \alpha = 0$ , if the fluid flux is outwards. This

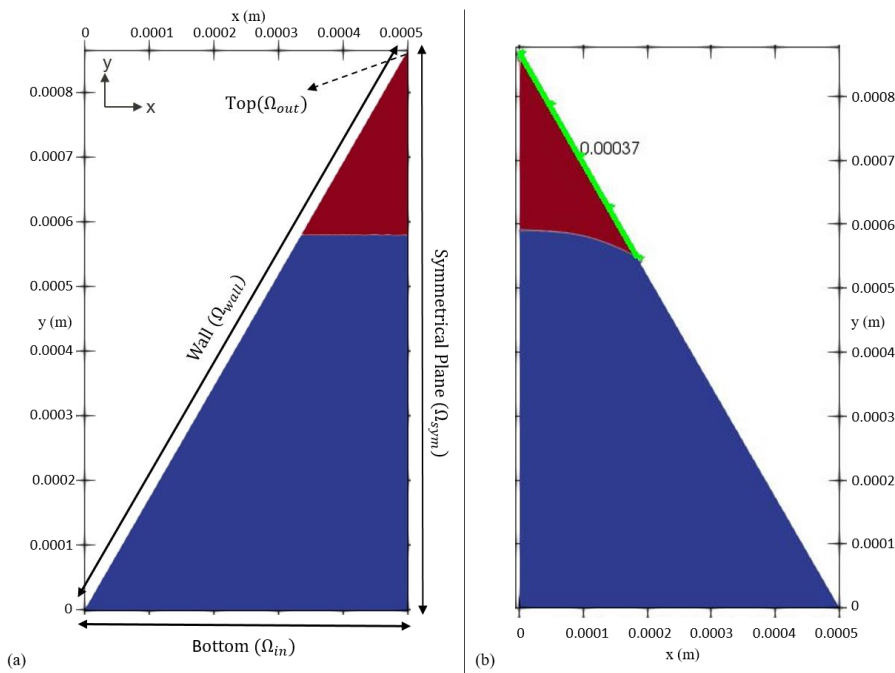
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374 ensured that appropriate fluids entered the domain from either inlet or outlet boundaries. The radius of  
 375 curvature can be also evaluated from the Young-Laplace equation ( $P_c = \frac{\sigma}{r}$ ). With a pressure difference of  
 376 255.33 ( $\text{kg m}^{-1} \text{s}^{-2}$ ) obtained from the last step and a surface tension of 0.07 ( $\text{kg s}^{-2}$ ), the radius of  
 377 curvature is calculated to be  $2.17\text{e-}4$  m. In a different approach, once the interface attains stationarity, we  
 378 calculated  $r$  for equation (24) as the reciprocal of the interface's mean curvature ( $2.77\text{e-}4$  m). For a  
 379 contact angle of  $10^\circ$  and a corner half-angle of  $30^\circ$ , the analytically calculated value for the length,  $b$ , of  
 380 the section in contact with the wetting phase is  $375 \mu\text{m}$ , while the numerical solution yields  $370 \mu\text{m}$ .  
 381 With a relative error of 1.21% this shows a reasonable match between numerical and analytical solutions  
 382 in modelling two-phase flow.

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384 **Figure 3:** Initial condition (a) versus final arrangement (b) of the two phases in the fluid configurations scenario. The blue color  
 385 indicates the non-wetting and the red color shows the wetting phases respectively. The dashed arrow shows the location of the  
 386 outlet, while the solid-line arrows depict the extent of others boundaries. Once equilibrium is reached (figure b), the curvature of  
 387 the interface corresponds to the force balance between pressure difference across the interface and the surface tension which can  
 388 be used to verify the model's sanity. The distance of the contact point (i.e. the point/line where all three phases – water, air and  
 389 solid – meets) from the corner vertex (highlighted as green), also provides another measure for checking the accuracy of the  
 390 numerical model.

391 3.2 Mass Transfer across the Fluid-Fluid Interface

392 Mass transfer of dissolved species between different phases, is particularly of importance for various  
393 biogeochemical processes in unsaturated subsurface environments ~~as e.g., oxygen or volatile organic~~  
394 ~~carbon compounds are found in the liquid and the gas phase, yet their transport and reaction conditions~~  
395 ~~differ highly between these two phases such as highly volatile contaminated water streams.~~

396 Model performance in simulating mass flux across the fluid-fluid interface is validated via a numerical  
397 experiment in which two immiscible stationary fluids (an aqueous -  $\alpha$  - and a gaseous -  $\beta$  - phase,  $\mathbf{u} = 0$  in  
398  $\Omega_\theta$ ) are horizontally (to remove buoyancy effects) residing on a one dimensional tube of 10 mm length  
399 with mesh size of 100  $\mu\text{m}$ . The general partial differential equation (PDE) of equation (12(12) takes the  
400 form of a simple diffusive transport as:

$$\frac{\partial C_{tr,i}}{\partial t} - D_{tr,i} \nabla(C_{tr,i}) = 0 \quad i = aq, gs$$

$$BC - 1: \quad C_{tr,aq} \times H = C_{tr,gs} \quad \text{in } \Omega_I \quad (25(25))$$

$$BC - 2: \quad D_{tr,aq} \frac{\partial C_{tr,aq}}{\partial x} = D_{tr,gs} \frac{\partial C_{tr,gs}}{\partial x} \quad \text{in } \Omega_I$$

401 with  $C_{tr}$  as the concentration of a volatile tracer and  $D_{tr,i}$  as the diffusivity of the tracer in phase  $i$ . Each  
402 phase is set to occupy half of the total volume (Figure 4(Figure 4-a). The system is initialized with a  
403 volatile chemical species of concentration of  $1 \text{ mol m}^{-3}$  in  $\Omega_{\theta,aq}$ , and  $0 \text{ mol m}^{-3}$  in  $\Omega_{\theta,gs}$ . At the inlet and  
404 the outlet boundary, tracer concentration equals that of the nearest solution such that, in short simulation  
405 time, it yields no concentration gradient into or out of the domain. The tendency of the dissolved chemical  
406 component to cross the fluid-fluid interface is expressed using a constant Henry coefficient. Tracer  
407 diffusivity is set to be  $1e-5 \text{ m}^2 \text{ s}^{-1}$  in both phases. The analytical solution for equation (25)(25) can be  
408 found in Bird (Bird, 2002).

409 Three scenarios with low, neutral, and high affinity of the volatile compound towards the gaseous phase  
410 are considered with corresponding Henry coefficients of 0.01 (low volatility, similar to naphthalene), 1  
411 (moderate volatility, e.g., vinyl chloride) and 100 respectively (high volatility, e.g. heptane). For a low  
412 value of  $H$  ( $H = 0.01$  - Figure 4(Figure 4-b) little (almost no) tracer is crossing the interface, while at  
413 neutral condition ( $H = 1$  - Figure 4(Figure 4-c), tracer diffusion is invariant to the phase it is occupying.  
414 Evidently for high values of  $H$  ( $H = 100$  - Figure 4(Figure 4-d), significant reduction in the tracer amount  
415 within the liquid phase, accompanied by a notable change in concentration across the interface, can be  
416 detected. This complies fully with the concentration jump due to for the such having highly volatile  
417 component in between the liquid and the gas phase. strong depletion of the tracer in aqueous phase as

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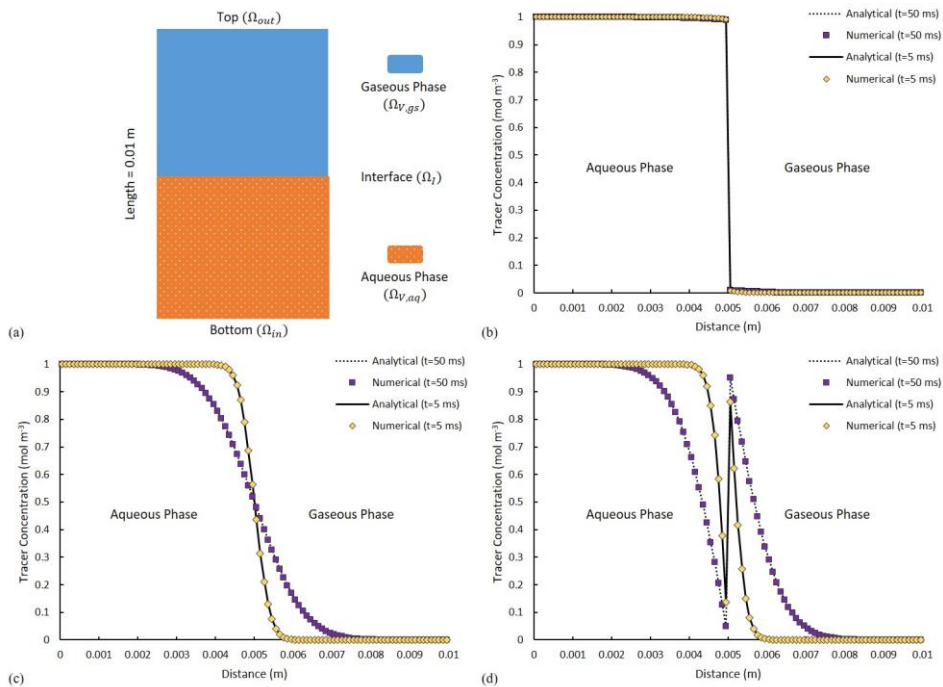
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418 well as the concentration jump across the interface are observed. The numerical results are ubiquitously  
 419 identical to the results of the analytical solution (Figure 4-b, c, d). The effect of grid  
 420 size/resolution is also investigated for this scenario. With 10 times higher grid resolution, the total CPU-  
 421 elapsed time is increased from ~650 seconds to ~3500 seconds.  
 422 The concentration profile remains unchanged, but, the average residuals of the numerical solution of the  
 423 concentration field, calculated at the end of the simulation, is increased from  $3.1e-10$  to  $6.5e-10$  (meaning  
 424 increasing resolution does not necessarily helps with numerical convergence).  
 425 Also, the average residuals of the numerical solution of the concentration field, calculated at the end of  
 426 the simulation, is increased from  $3.1e-10$  to  $6.5e-10$  (meaning increasing resolution does not necessarily  
 427 helps with numerical convergence).

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428  
 429 **Figure 4:** (a) A schematic of the fluid distributions at initial condition. The solid and dotted lines show the analytical solutions  
 430 with purple and yellow squares depicting the numerical solutions. Comparison of the analytical and numerical solutions of tracer  
 431 distribution at two distinct time points of  $t_1=5ms$  and  $t_2=50ms$  for (b)  $H = 0.01$ , (c)  $H = 1$ , and (d)  $H = 100$ .

432

433 3.3 Microbial Growth and Reactive Transport

434 Our modelling framework can parameterize any type of reactions, we put the main focus in this  
 435 subsection on microbially driven redox transformations (i.e. a type of reactions commonly encountered in  
 436 soils and other porous media environments) and on the implementation of the corresponding  
 437 mathematical formulation. To validate our model with a scenario in which bacterial biomass is allowed to  
 438 evolve (i.e. to grow and to decay) we adapted a conceptual biodegradation scenario from (Cirpka and  
 439 Valocchi, 2007b) in which a fully-water-saturated, two-dimensional domain-channel is subjected to a  
 440 constant flux of two different components; ED (electron donor – e.g. hydrocarbon) and EA (electron  
 441 acceptor – e.g. oxygen). The imposed uniform flow field is assumed to be constant over time and has only  
 442 the x-component. The bacteria residing in the channel, facilitate the reaction between ED and EA, which  
 443 can be written in an abstract form as  $f_a ED + f_b EA \xrightarrow{\text{biomass}} f_c Prod$ , where *biomass* is the microbial  
 444 biomass, *Prod* is the product(s) (e.g. metabolites such as carbon dioxide) and  $f_a, f_b$  and  $f_c$  are  
 445 stoichiometric coefficients. Assuming a double-Monod kinetics for expressing microbial growth and the  
 446 microbially driven reaction rates, as well as assuming none of the reactants nor products are involved in  
 447 secondary reactions, the ADRE (eq. (12)+(2)) for each chemical species can then be written as:

$$\begin{aligned} \frac{\partial C_{ED}}{\partial t} + \mathbf{u} \frac{\partial C_{ED}}{\partial x} - D_t \frac{\partial^2 C_{ED}}{\partial y^2} &= - \frac{C_{ED}}{C_{ED} + K_{ED}} \frac{C_{EA}}{C_{EA} + K_{EA}} \frac{\mu_{max}}{Y} f_a C_{bio} & \text{in } \Omega_\theta \\ \frac{\partial C_{EA}}{\partial t} + \mathbf{u} \frac{\partial C_{EA}}{\partial x} - D_t \frac{\partial^2 C_{EA}}{\partial y^2} &= - \frac{C_{ED}}{C_{ED} + K_{ED}} \frac{C_{EA}}{C_{EA} + K_{EA}} \frac{\mu_{max}}{Y} f_b C_{bio} & \text{in } \Omega_\theta \\ \frac{\partial C_{Prod}}{\partial t} + \mathbf{u} \frac{\partial C_{Prod}}{\partial x} - D_t \frac{\partial^2 C_{Prod}}{\partial y^2} &= \frac{C_{ED}}{C_{ED} + K_{ED}} \frac{C_{EA}}{C_{EA} + K_{EA}} \frac{\mu_{max}}{Y} f_c C_{bio} & \text{in } \Omega_\theta \\ \frac{\partial C_{bio}}{\partial t} &= \frac{C_{ED}}{C_{ED} + K_{ED}} \frac{C_{EA}}{C_{EA} + K_{EA}} \mu_{max} C_{bio} - \lambda C_{bio} & \text{in } \Omega_\theta \end{aligned} \quad (2626)$$

448 where  $\mathbf{u}$  is the velocity (which has only a constant  $x$ -component),  $D_t$  is the transverse dispersivity,  
 449  $C_{ED}, C_{EA}, C_{Met}$  and  $C_{bio}$  are concentrations of *ED, EA, Prod* and *biomass* respectively;  $K_{ED}$  and  $K_{EA}$  are  
 450 half saturation constants for respective compounds in the biomass growth term,  $Y$  is the yield coefficient,  
 451  $\mu_{max}$  is the maximum bacterial growth rate, and  $\lambda$  is the bacterial decay rate. Using these equations  
 452 Cirpka and Valocchi (2007a) developed an analytical solution for steady-state conditions, which in the  
 453 version of Cirpka and Valocchi (2009) is used as reference for the numerical results.

454 The numerical experiment is designed to have ED and EA, occupying 25% and 75% of the inlet  
 455 respectively, and, simultaneously, invading the domain under a constant uniform velocity field, with  
 456 concentration of  $C_{ED}^{inlet}$  and  $C_{EA}^{inlet}$ . In a real-world scenario, this can be seen as a plume of a contaminant

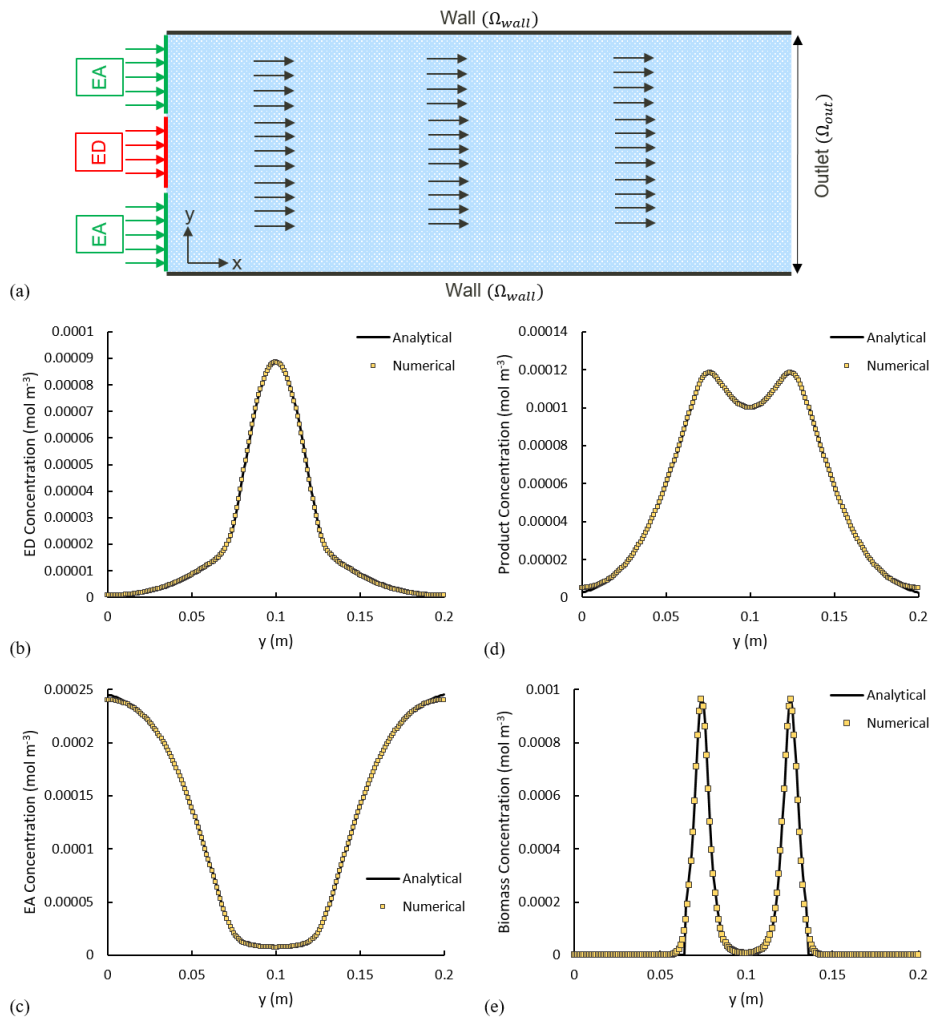
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457 (i.e. a hydrocarbon as *ED*) being carried into the domain within an oxygenated stream and essentially we  
458 are interested in knowing the final concentration/distribution of all bio-chemical species within the  
459 domain. The parameters used in this scenario are summerized in Table 1~~Table 1~~. Transient reactive  
460 transport ~~simluations~~simulations are performed until a steady state is achieved. For validation, we analyze  
461 all concentration profiles along the *y*-axis at a fixed distance of  $x = 2$  m and compare them with the  
462 analytical solutions. The analytical and numerical results show an almost perfect agreement (Figure  
463 5~~Figure 5~~, b-e).

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465 **Figure 5:** (a) Model set up. The arrows show the direction of the flow field. Solid lines show the analytical solution and the  
 466 yellow squares illustrate the numerical results. (b-e) Comparison of the analytical and numerical solutions  $x=2m$  for  
 467 concentration of (b) *Electron Donor*, (c) *Electron Acceptor*, (d) *Product*, and (e) *Biomass*.

468

469

470

Table 1: Parameter values used for simulating microbial growth.

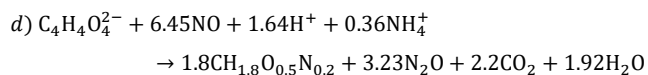
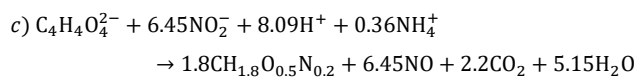
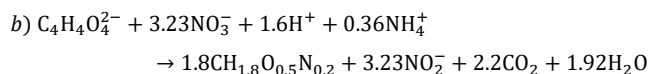
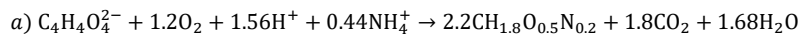
Parameter	Value	Unit	Parameter	Value	Unit
$u$	100	$cm\ d^{-1}$	$\mu_{max}$	1	$d^{-1}$
Domain width	20	$cm$	Domain length	500	$cm$
$D_t$	2.5	$cm^2\ d^{-1}$	$\lambda$	0.1	$d^{-1}$
$f_a, f_b, f_c$	1	-	$K_{ED}$	8.33e-2	$mM$
$C_{ED}^{inlet}$	0.33	$mM$	$K_{EA}$	3.13e-2	$mM$
$C_{EA}^{inlet}$	0.25	$mM$	Y	1	-
Mesh size	0.2	mm			

471

## 472 3.4 Theme: Demonstrating Model Capabilities

473 The scenarios described above are designed to serve as the sole purpose of creating a baseline for  
474 validating the numerical toolbox – simple enough where analytical solutions could exist. Unlike the  
475 simplicity introduced in previous sections, simulating soil processes with all of the complexities, though,  
476 would require having all the modelling elements to be present. We thus present here a scenario with an  
477 unsaturated soil hosting the facultative anaerobic bacteria *Agrobacterium tumefaciens*, which performs  
478 aerobic respiration under oxic condition, but switches to denitrification using nitrate, nitrite or nitric oxide  
479 under anoxic condition (Kampschreur et al., 2012). This example allows us to show our model  
480 capabilities, as it involves 1) the actual micro structure of the soil, 2) unsaturated conditions, and 3) an  
481 enzymatic reaction network with limiting/inhibition terms. The microstructure is obtained via  
482 subsampling from a larger  $\mu$ -CT image with voxel size of 6 micron (see Supplementary Information). A  
483 two-phase simulation is then performed on the voxelized subsample to obtain the fluxes and phase  
484 distribution of air and water within the pore space. For this, the entire domain is initially filled with water  
485 and subject to injecting air from the top boundary with constant flux of  $0.013\ ml\ h^{-1}$ . An important note to  
486 make here is with a relatively high influx, advection transport acts as the bottleneck for numerical time  
487 steps. Hence, reactions are performed at a quite slower pace (i.e. larger time steps roughly estimated  
488 around 10 hrs). This separation of processes helps improve the overall run-time of the simulations.  
489 Generally, the time step sizes are automatically enforced by the Courant number from the transient  
490 advective-diffusive transport equation (with order of  $10^{-5}$  seconds). The biomass is assumed to be non-  
491 motile meaning it sticks to the solid surface and shows no planktonic behavior. ~~to be~~ Fluids are  
492 allowed to leave the domain from the bottom part (kept at atmospheric pressure) while all the remaining  
493 sides are set to be impermeable walls. Once fluid configurations in the domain are stationary, their

494 distribution along with the velocity profile are used as basis for the reactive transport simulations (phase  
 495 distributions can be found in the Supplementary Information). Using succinate ( $C_4H_4O_4^{2-}$ ) as organic  
 496 carbon substrate to be degraded, a metabolic reaction network is constructed with four microbial  
 497 degradation pathways each following Monod-type kinetics: 1) aerobic respiration with a nitric oxide ( $NO$ )  
 498 inhibitory term, 2) nitrate ( $NO_3^-$ ) reduction, 3) nitrite ( $NO_2^-$ ) reduction and 4)  $NO$  reduction, with having  
 499 oxygen ( $O_2$ ) as inhibitory element for all denitrification conversions (eq. (27)(27)). Also three additional  
 500 equations are considered for the synthesis of the three different enzymes required for degradation  
 501 processes (eq. (28)(28)). We consider only one single strain of bacteria (*Agrobacterium tumefaciens*) which  
 502 has the benefits of performing both aerobic respiration and denitrification. Bacteria are considered to be  
 503 non-motile with an initial concentration of  $0.25 \text{ mol m}^{-2}$  and uniformly covering the entire grain surface  
 504 area. Succinate has its initial concentration in the aqueous phase set at  $0.2 \text{ mM}$  ( $0 \text{ mM}$  in the gaseous  
 505 phase), while all other species have their initial concentrations of  $0 \text{ mM}$  in both aqueous and gaseous  
 506 phases. Boundary condition for all concentration fields on all boundaries is set to zero gradient except for  
 507 the inlet boundary (fully saturated with air) - where for oxygen it is set to  $0.03567 \text{ mM}$ , and for all others  
 508 is set to  $0 \text{ mM}$ . In order to avoid depletion of the nitrate in the system, a nitrate concentration of  $0.1 \mu\text{M}$   
 509 (as initial condition) is provided. The complete reaction network can be written as follows (Kampschreur  
 510 et al., 2012):



511 Several assumptions are made for preparing the kinetics of the reactions: 1) reaction rates are limited by  
 512 the maximum specific uptake rate of succinate and are hence independent of its concentration (Beun et  
 513 al., 2000), 2) sufficient amount of buffer is added to the solution to keep the pH level constant, 3) three  
 514 nitrogen reductase enzymes ( $\xi_{sat,NOR}$  for  $NO$  reduction,  $\xi_{sat,NIR}$  for nitrite reduction and  $\xi_{sat,NAP}$  for  
 515 nitrate reduction) can have saturation values varying between 0 (i.e. non-existing) and 1 in a bacterial cell,  
 516 and 4) inhibitory oxygen limits the reduction of  $NO$ ,  $NO_2^-$  and  $NO_3^-$ . Reaction rates are designed to have a  
 517 dependency on the enzymes' level and biomass concentration with proper limiting/inhibiting terms.

518 Equation (12)(12) is used to describe the evolution of each biochemical species. The final system of  
 519 advective-diffusive-reactive equations is adapted from (Kampschreur et al., 2012):

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(27)(27)

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$$\frac{\partial C_i}{\partial t} + \nabla \cdot (C_i \mathbf{u}) = -\nabla \cdot \left( D_i \nabla C_i - D_i \frac{1 - H_i}{\alpha + (1 - \alpha) H_i} C_i \nabla \alpha \right) + R_i,$$

$$i = \text{suc}, O_2, NO_3^-, NO_2^-, NO, N_2O$$

in  $\Omega_\theta$

$$R_{\text{suc}} = -(r_{\text{suc},O_2} + r_{\text{suc},NAP} + r_{\text{suc},NIR} + r_{\text{suc},NOR})$$

$$R_{O_2} = -1.2r_{\text{suc},O_2}$$

$$R_{NO_3^-} = -3.23r_{\text{suc},NAP}$$

$$R_{NO_2^-} = 3.23r_{\text{suc},NAP} - 6.45r_{\text{suc},NIR}$$

$$R_{NO} = 6.45r_{\text{suc},NIR} - 6.45r_{\text{suc},NOR}$$

$$R_{N_2O} = 3.23r_{\text{suc},NOR}$$

$$r_{\text{suc},O_2} = \mu_{\text{max}} C_{\text{bio}} \frac{C_{O_2}}{K_{O_2} \left( 1 + \frac{C_{NO}}{K_{I,NO,O_2}} \right) + C_{O_2}}$$

$$r_{\text{suc},NAP} = \mu_{\text{max}} C_{\text{bio}} \xi_{\text{sat},NAP} \frac{C_{NO_3}}{K_{NO_3} + C_{NO_3}} \frac{K_{I,O_2,NAP}^{nNAP}}{K_{I,O_2,NAP}^{nNAP} + C_{O_2}^{nNAP}}$$

$$r_{\text{suc},NIR} = \mu_{\text{max}} C_{\text{bio}} \xi_{\text{sat},NIR} \frac{C_{NO_2}}{K_{NO_2} + C_{NO_2}} \frac{K_{I,O_2,NIR}^{nNIR}}{K_{I,O_2,NIR}^{nNIR} + C_{O_2}^{nNIR}}$$

$$r_{\text{suc},NOR} = \mu_{\text{max}} C_{\text{bio}} \xi_{\text{sat},NOR} \frac{C_{NO}^2}{\left[ C_{NO} \left( 1 + \frac{C_{NO}}{K_{I,NO}} \right) + K_{NO} \right]^2} \frac{K_{I,O_2,NOR}}{K_{I,O_2,NOR} + C_{O_2}}$$

$$\frac{d\xi_{\text{sat},NAP}}{dt} = v_{m,NAP} \frac{C_{NO_3}}{K_{NO_3,NAP} + C_{NO_3}} \frac{K_{I,O_2,NAP}}{K_{I,O_2,NAP} + C_{O_2}} (1 - \xi_{\text{sat},NAP})$$

$$\frac{d\xi_{\text{sat},NIR}}{dt} = v_{m,NIR} \frac{C_{NO_2}}{K_{NO_2,NIR} + C_{NO_2}} (1 - \xi_{\text{sat},NIR})$$

$$\frac{d\xi_{\text{sat},NOR}}{dt} = v_{m,NOR} \frac{C_{NO}}{K_{NO,NOR} + C_{NO}} (1 - \xi_{\text{sat},NOR})$$

$$\frac{dC_{\text{bio}}}{dt} = 2.2r_{\text{suc},O_2} + 1.8(r_{\text{suc},NAP} + r_{\text{suc},NIR} + r_{\text{suc},NOR})$$

(2828)

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520 The full list of modelling parameters used for this study can be found in [the Table S.2 \(see Supplementary](#)

521 Information).

522 Reactive transport simulations were performed until a quasi-steady state was achieved. This was

523 characterized by all chemical species concentrations reaching a steady-state as determined by the

524 degradation activity of the given distribution of microorganisms. Since microbial growth takes place at

525 much larger time scales than the pore-scale transport processes no significant growth takes place during

526 the simulated time period [and shown results are nearly identical to the initial conditions](#). Simulation

527 results show that the presence of air in this two-phase system affects the distribution of biochemical  
528 species. Air, as the non-wetting phase, occupies the central part of the pore space while the aqueous phase  
529 is expected to cover the corners and crevices (Figure 6, a). For oxygen with  $H_{O_2} = 31$  a higher  
530 concentration is observed in the air compared to that of the adjacent aqueous phase (Figure 6, d).  
531 An analysis of how the volatility of a tracer compound may affect its residence time in the porous  
532 medium is given in the Supplementary Information. Since the biomass is only present on the grain  
533 surfaces (Figure 6, b), oxygen, nitrate and succinate deplete as the microbially-mediated reactions  
534 only at these micro-locations. Fresh oxygen and nitrate thus need to diffuse from the bulk (either from the  
535 aqueous phase or the air) to the reactive sites. The regions with high (i.e. not degraded) succinate  
536 concentrations are compatible with low concentration regions of oxygen and nitrate, i.e. the reactions are  
537 limited by the bioavailable oxygen and nitrate (Figure 6, b-e). Finally, all three enzymes have an  
538 increased abundance in anaerobic regions with an active biomass (saturation map of nitrate reductase  
539 enzyme is shown in Figure 6, e). While the saturation of nitrate reductase enzyme grows linearly  
540 with time (until 0.25 s), the rate at which the nitrite and NO reductase enzymes ( $\xi_{sat,NIR}$  and  $\xi_{sat,NOR}$   
541 respectively) ~~growth-grow~~ is rather slow for the very beginning of the simulation (until ~0.2 s), but it  
542 surges exponentially afterward. A spatially integrated assessment of the degradation processes showed  
543 that for the presented example 99% of the total succinate degradation is attributed to aerobic respiration  
544 while a trivial amount is attributed to the three anaerobic processes (nitrate reduction, nitrite reduction and  
545 NO reductions).

546 The presented results highlight the ability of the model to combine a high-resolution simulation of multi-  
547 phase flow and transport processes with the simulation of complex biogeochemical processes. This allows  
548 for a realistic simulation of the ~~pore-scale~~~~micro-scale~~ distribution of reactive processes and for the  
549 derivation of an accurate aggregated description of these processes.

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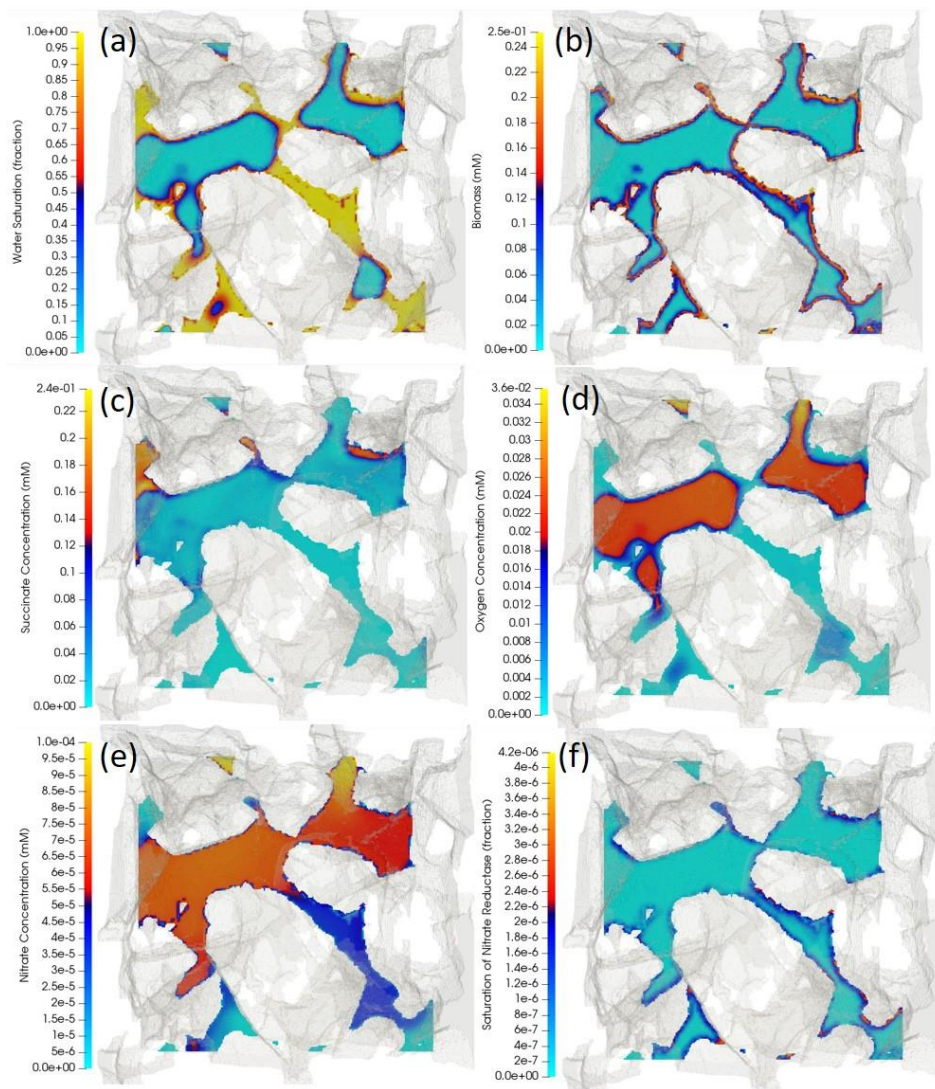
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550

551 **Figure 6** Cross-sectional view of the three dimensional porous medium. The cutting plane is arbitrary cutting through the middle  
 552 of the porous structure, meaning at some locations, the phases are continuous perpendicular to the plane. The opaque grayish  
 553 background represents the 3D porous structure that is extracted and digitized from a  $\mu$ -CT image. The colored surfaces are  
 554 obtained by running a cutting plane through the middle of the sample and perpendicular to the z-axis. The distribution of (a)  
 555 water-content fraction (i.e. water volume fraction), (b) biomass, (c) succinate, (d) oxygen, (e) nitrate and (f) nitrate reductase  
 556 enzyme are respectively depicted with having yellow color indicating highest value and light blue as the lowest value. With air as  
 557 the non-wetting phase, it is expected to fill in the middle of the pore space where capillary pressure is lower while water, as the  
 558 wetting phase, is expected to occupy the corners (figure a). A high volatility constant for oxygen enforces to have higher  
 559 concentrations of oxygen in the air compared to that of aqueous phase adjacent to the water-air interface. Cross-sectional view of

560 the three dimensional porous medium. The opaque grayish background represents the 3D porous structure that is extracted and  
561 digitized from a  $\mu$ -CT image. The colored surfaces are obtained by running a cutting plane through the middle of the sample and  
562 perpendicular to the z axis. The distribution of (a) water saturation (i.e. water volume fraction), (b) biomass, (c) succinate, (d)  
563 oxygen, (e) nitrate and (f) nitrate reductase enzyme are respectively depicted with having yellow color indicating highest value  
564 and light blue as the lowest value. With air as the non-wetting phase, it is expected to fill in the middle of the pore space where  
565 capillary pressure is lower while water, as the wetting phase, is expected to occupy the corners (figure a). A high volatility  
566 constant for oxygen enforces to have higher concentrations of oxygen in the air compared to that of aqueous phase adjacent to the  
567 water-air interface.

568 As it can be seen from [Figure 6](#), our model can be used (among other options) to identify clusters  
569 in which succinate is most and least depleted. This would ease the process of analyzing the results by  
570 isolating the parameters that are boosting/limiting the degradation of the carbon source. 3D visualization  
571 of the oxygen and succinate distributions can be found in (Golparvar, 2022).

#### 572 4. Conclusion and future remarks

573 In this paper, we have presented a newly developed modelling framework for simulating reactive  
574 transport processes in real porous soil structures obtained from  $\mu$ -CT images under unsaturated  
575 conditions. The successful application of various benchmark test showed the model's accuracy in the  
576 simulation of 1) the movement of water and air phase in variably saturated conditions via the enhanced  
577 algebraic Volume of Fluid method (Raeni et al., 2012) coupled with the Navier Stokes equation, 2) the  
578 transport of different species in both phases by the full advective-diffusive transport equation, and finally  
579 3) using the operator splitting technique, an arbitrary set of biogeochemical reactions solved externally by  
580 the Biogeochemical Reaction Network Simulator and communicated back into the main solver.

581 The presented model provides a novel and unique combination of pore-scale simulations of two-phase  
582 flow, transport of dissolved and volatile species and their reactive transformations. This makes it an  
583 accurate and powerful tool for the simulation of soil systems or other unsaturated porous media and of the  
584 reactive transport processes therein. While developed with the aim for simulating biogeochemical  
585 processes in soils the model is equally applicable for simulating other abiotic reactive processes coupled  
586 to the dynamics of flow and transport in variably saturated pore structures of arbitrary geometry. Our  
587 modelling framework is properly designed for simulating biogeochemical processes such as carbon-  
588 nitrogen-sulfur-phosphorus cycles in soil as well as mixing and migration of contaminants in both  
589 unsaturated soil and water aquifers. It comes with the benefit of explicit recognition of the soil structure  
590 (i.e. using the 3D structure as close to the original shape as possible with the least amount of  
591 simplifications/modifications), phase dynamics/distributions and the capability of designing the complete  
592 redox reactions necessary for a given process in a straightforward fashion. It is best suitable for running  
593 pre-pilot tests as feasibility scenarios where the stakes for the success of the project are high. Also, our  
594 model provides the best tool for designing hypothetical experiments that are hard (if not impossible) to  
595 implement experimentally (e.g. a specific distribution of biomass/reactants within the domain, or

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596 variation of specific properties of reactive compound and/or the porous matrix). Furthermore, the high  
597 resolution modelling results provided by this model support the upscaling of reactive-transport process  
598 description from the pore to the continuum scale and from the process to the observation scale,  
599 respectively.

600 Although the current version of our numerical model is already covering a wide range of bio-physio-  
601 chemical properties of the soil constituents, for having more realistic representation of multiphase,  
602 multicomponent reactive transport in partially saturated porous media, few more factors still might be  
603 considered in future developments of the model: 1) shrinkage/expanding of the air/aqueous phase due to  
604 mass transfer of chemical species across fluid-fluid interface, 2) accounting for gas compressibility by  
605 adding an equation of state for tracking changes of air volume/density under flowing condition, 3)  
606 translating accumulated biomass on the grain surfaces into new flow-resistance components which would  
607 potentially change the velocity streamlines (i.e. bioclogging), 4) changes of the grain surface structures  
608 and of the associated solid-liquid interface due to mineral precipitation/dilution or due to  
609 accumulation/depletion of solid organic material, ~~and~~ 5) chemotactic behavior of the microbial species,  
610 and 6) osmotic forces and electro-migration. Due to its modular structure, such features can be relatively  
611 easily included into future upgrades of our model. -

## 612 **5. Acknowledgements**

613 This work was funded by the Helmholtz Association via the integrated project “Controlling Chemicals  
614 Fate.”

## 615 **6. Code and data availability**

616 The source codes, benchmark and demonstration cases along with instruction for installing and running  
617 each case that are presented in this paper, are archived at <https://github.com/amirgolp/P3D-BRNS> DOI:  
618 [10.5281/zenodo.6301317](https://doi.org/10.5281/zenodo.6301317).

## 619 **7. Author contributions**

620 AR was responsible for model/software curation, validation and visualization. Conceptualization and  
621 methodology development were managed by AR, MK, and MT. Writing the original manuscript was  
622 handled by AR while all authors contributed to the revision and curation of the final draft. The entire  
623 work is supervised by MT.

## 624 **8. Competing interests**

625 The authors declare that they have no conflict of interest.

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