



1	AttentionFire_v1.0: interpretable machine learning fire model for burned area
2	predictions over tropics
3	Fa Li ^{1,2} , Qing Zhu ^{1,*} , William J. Riley ¹ , Lei Zhao ³ , Li Xu ⁴ , Kunxiaojia Yuan ^{1,2} , Min
4	Chen ⁵ , Huayi Wu ² , Zhipeng Gui ⁶ , Jianya Gong ⁶ , James T. Randerson ⁴
5	¹ Climate and Ecosystem Sciences Division, Climate Sciences Department, Lawrence
6	Berkeley National Laboratory, Berkeley, CA, USA
7	² State Key Laboratory of Information Engineering in Surveying, Mapping and
8	Remote Sensing, Wuhan University, Wuhan, China
9	³ Department of Civil and Environmental Engineering, University of Illinois Urbana-
10	Champaign, Champaign, IL, USA
11	⁴ Department of Earth System Science, University of California Irvine, Irvine, CA,
12	USA
13	⁵ Department of Forest and Wildlife Ecology, University of Wisconsin-Madison,
14	Madison, WI, USA
15	⁶ School of Remote Sensing and Information Engineering, Wuhan University, Wuhan,
16	China
17	*Correspondence to Qing Zhu (qzhu@lbl.gov)
18	
19	Abstract
20	African and South American (ASA) wildfires account for more than 70% of global
21	burned areas and have strong connection to local climate for sub-seasonal to seasonal
22	wildfire dynamics. However, representation of the wildfire-climate relationship





remains challenging, due to spatiotemporally heterogenous responses of wildfires to 23 24 climate variability and human influences. Here, we developed an interpretable Machine Learning (ML) fire model (AttentionFire v1.0) to resolve the complex spatial-25 26 heterogenous and time-lagged controls from climate on burned area and to better 27 predict burned areas over ASA regions. Our ML fire model substantially improved predictability of burned area for both spatial and temporal dynamics compared with 28 29 five commonly used machine learning models. More importantly, the model revealed 30 strong time-lagged control from climate wetness on the burned areas. The model also 31 predicted that under a high emission future climate scenario, the recently observed declines in burned area will reverse in South America in the near future due to climate 32 changes. Our study provides reliable and interpretable fire model and highlights the 33 importance of lagged wildfire-climate relationships in historical and future predictions. 34

35

36 1. Introduction

37 Wildfires modify land surface characteristics, such as vegetation composition, soil 38 carbon, surface runoff, and albedo, with significant consequences for regional carbon, 39 water, and energy cycles [Benavides-Solorio and MacDonald, 2001; Randerson et al., 2006; Shvetsov et al., 2019]. Over African and South American (ASA) regions, where 40 more than 70% of global burned area occurs, wildfires emit ~1.4 PgC yr⁻¹ and dust and 41 42 aerosols that can alter regional climate through radiative processes [Etminan et al., 2016; Ramanathan et al., 2001; Werf et al., 2017]. While greenhouse gas emissions contribute 43 to climate change, other toxic species and airborne particulate matter from wildfires 44





45	lead to substantial health hazards, including elevated premature mortality [Knorr et al.,
46	2017; Lelieveld et al., 2015]. In particular, wildfire particulate matter emissions across
47	tropical regions have exceeded current anthropogenic sources and are predicted to
48	dominate future regional emissions [Knorr et al., 2017].
49	Although total tropical wildfire burned area has declined over the past few decades
50	due to climate change and human activities [Andela et al., 2017; Andela and Van Der
51	Werf, 2014] (e.g., from increases in population density, cropland fraction, and livestock
52	density), wildfire still plays a significant role in mediating surface climate [Xu et al.,
53	2020], biogeochemical cycles, and human health [Andela et al., 2017]. Further, 21st
54	century projections of increases in temperature, regional drought [Dai, 2013; Taufik et
55	al., 2017], and precipitation variations may outweigh these direct human impacts and
56	result in unprecedentedly fire-prone environments over a large fraction of Africa
57	[Andela and Van Der Werf, 2014; Archibald et al., 2009; Van Der Werf et al., 2008] and
58	South America [Malhi et al., 2008; Pechony and Shindell, 2010]. These factors
59	highlight the need for better understand, predict, and management of these critical fire
60	regions to minimize economic losses, human health hazards, and natural ecosystem
61	degradation. Therefore, improved understanding and accurate prediction of wildfire
62	activity is increasingly important for effective fire management and sustainable
63	decision-making.

Climate is acknowledged one of the most dominant controllers on ASA wildfires 64 [Andela et al., 2017; Chen et al., 2011]. For example, precipitation variations contribute 65 substantially to burned area patterns in southern and northern Africa [Andela and Van 66





67	Der Werf, 2014; Archibald et al., 2009], and are also closely linked to wildfire
68	spatiotemporal dynamics in south America [Chen et al., 2011; Malhi et al., 2008; Van
69	Der Werf et al., 2008]. More importantly, the strong controls from climate on wildfires
70	often show time-lags and the time-delay can be up to multiple months [Andela and Van
71	Der Werf, 2014; Van Der Werf et al., 2008], which enables wildfire predictions ahead
72	of fire season [Chen et al., 2016; Chen et al., 2020; Chen et al., 2011; Turco et al.,
73	2018]. The spatiotemporal responses of wildfires to climate changes are complicated
74	by non-linear interactions among climate, vegetation, and human activities [Andela et
75	al., 2017; Van Der Werf et al., 2008]. In more xeric subtropical regions, increasing
76	precipitation during the wet season can be the dominant controller on increasing
77	wildfire during the following dry season (through regulation of fuel availability and
78	fuel spatial structures) [Archibald et al., 2009; Littell et al., 2009; Van Der Werf et al.,
79	2008]. In contrast, increasing precipitation in more mesic regions results in excessive
80	fuel moisture, thereby becoming the main limitation of dry-season wildfires (i.e.,
81	opposing fire trends are observed with increasing precipitation in northern and southern
82	Africa) [Andela and Van Der Werf, 2014; Van Der Werf et al., 2008]. In addition to
83	natural processes, human activities are primary ignition sources and have shaped fire
84	patterns in the ASA regions [Andela et al., 2017; Aragao et al., 2008; Archibald et al.,
85	2009]. Fire-use types driven by local socio-economic conditions and fire management
86	policies may also affect the fire-climate relationships [Andela et al., 2017]. Therefore,
87	strong climate controls from wet season to dry season need to be considered along with





fuel distributions and human activities for continental fire predictions under climate

Accurate predictive modeling of wildfire with skillful representation of how 90 environmental and anthropogenic factors modulate the burned area is still challenging. 91 State-of-the-art process-based fire models (e.g., the Fire Model Intercomparison Project 92 [Rabin et al., 2017]) have reasonably simulated the spatial distribution of burned areas. 93 94 However, they generally do not accurately capture burned area seasonal variation and 95 inter-annual trends and variability [Andela et al., 2017]. Improving predictability and 96 reducing uncertainties of process-based models require more sophisticated representation of fire processes and parameterization, which remain a long-term 97 challenge [Bowman et al., 2009; Hantson et al., 2016; Teckentrup et al., 2019]. In 98 response to this challenge, data-driven statistical or Machine Learning (ML) 99 100 approaches have been developed and demonstrated to effectively capture wildfire severity and burned area dynamics [Archibald et al., 2009; Chen et al., 2020; Chen et 101 al., 2011; Zhou et al., 2020]. However, the spatially heterogenous, non-linear, and time-102 103 lagged controls have been oversimplified (e.g., using linear models or only considering climate variables at specific time lags or seasons [Archibald et al., 2009; Chen et al., 104 2016; Chen et al., 2020; Chen et al., 2011; Gray et al., 2018a]) or have been black 105 boxed, impeding an interpretable and reliable way to understand the critical 106 107 spatiotemporal processes from wet season to dry season [Jain et al., 2020; Reichstein et al., 2019]. 108

⁸⁹ change.





109	In this work, we developed a wildfire model (AttentionFire) leveraging on an
110	interpretable Long-Short-Term-Memory framework to predict wildfire burned areas
111	over Northern Hemisphere Africa (NHAF), Southern Hemisphere Africa (SHAF), and
112	Southern Hemisphere South America (SHSA) [Giglio et al., 2013]. We also focused on
113	using the AttentionFire model to explore the dependency of simulated burned area on
114	different drivers from wet season to dry season across different gridcells. We assessed
115	model predictability with observed burned area from Global Fire Emission Database
116	(GFED) and compared with five other machine learning based fire models.

117

118 **2. Methods**

119 **2.1 AttentionFire model**

120 The AttentionFire model is based on an interpretable attention-augmented LSTM 121 [Guo et al., 2019; Li et al., 2020; Liang et al., 2018; Oin et al., 2017; Vaswani et al., 2017] framework. An naïve LSTM has shown advantages in capturing short- and long-122 term dependencies in input time series [Hochreiter and Schmidhuber, 1997]. However, 123 124 LSTM cannot explicitly and dynamically select important drivers from multiple driving time series to make predictions [Guo et al., 2019; Li et al., 2020; Liang et al., 2018; 125 Qin et al., 2017; Vaswani et al., 2017]. Further, LSTM works as a black-box, lacking 126 interpretability to identify the relative importance of each driver across different time 127 128 steps [Guo et al., 2019; Li et al., 2020; Liang et al., 2018]. Attention mechanisms overcome these challenges by adaptively assigning larger weights to more important 129 drivers and time steps [Liang et al., 2018; Vaswani et al., 2017]. Here we use attention 130





131 mechanism to explicitly capture controlling factors of fire predictions with various







Fig. 1: An illustrative workflow for AttentionFire_v1.0 model prediction. Four kinds of drivers are considered: ignition related, suppression related, fuel, and climate. The temporal attention is used to identify important time steps for each kind of driver, while the variable attention is used to identify important drivers for final burned area prediction.

139

Given four categories of time series, $X = (X^1, X^s, X^f, X^c)^T$, where *T* is the length of time series, we use $X^i = (x_1^i, x_2^i, ..., x_T^i)^T \in \mathbb{R}^T$, where $1 \le i \le n$, to denote the *i*th time series, and use $X_t = (x_t^1, x_t^2, ..., x_t^n)^T \in \mathbb{R}^n$, where $1 \le t \le T$, to represent the vector at time step *t*. x_t^l , x_t^s , x_t^f , and x_t^c represent the variables of ignition (*e.g.*,





144 population density), suppression (*e.g.*, road network density), fuel availability (*e.g.*, 145 living biomass), and climate (*e.g.*, precipitation) at time step t. The AttentionFire 146 model aims to learn a nonlinear function F to map the n time series to the observed 147 burned area Y_{T+1} at time step T + 1:

$$\widehat{Y}_{T+1} = F(X^{\mathrm{l}}, X^{\mathrm{s}}, X^{\mathrm{f}}, X^{\mathrm{c}})^{T}$$

$$\tag{1}$$

148 Where \hat{Y}_{T+1} is the predicted burned area at time step T + 1.

First, the model iteratively transforms the *i*-th driving variable at time step *t* to a hidden state vector h_t^i , where $1 \le t \le T$ and $1 \le i \le n$ through LSTM gate mechanisms [*Guo et al.*, 2019; *Li et al.*, 2020; *Liang et al.*, 2018; *Qin et al.*, 2017]. Second, as the importance of each time step varies, temporal attention is applied to h_t^i to calculate its corresponding weight or importance w_t^i . Third, the weighted summation h_{sum}^i of h_t^i is obtained to represent the summarized information for the *i*-th driving variable:

$$w_t^i = f_{attn}(h_t^i)$$

$$h_{sum}^i = \sum_{t=1}^T w_t^i h_t^i$$
(2)

156 Where $h_t^i \in \mathbb{R}^m$ is the hidden state vector of the *i*-th driving series at time step *t*, that 157 stores the summary of the past input sequence [*Hochreiter and Schmidhuber*, 1997]. 158 w_t^i is the calculated weight for the *i*-th driver at time step *t* through attention function 159 f_{attn} :

$$w_t^{i'} = \tanh\left(W_p h_t^i\right) \tag{3}$$

8





$$w_t^i = \frac{e^{w_t^{i'}}}{\sum_{j=1}^T e^{w_t^{j'}}}$$

160 where $W_p \in \mathbb{R}^{1 \times m}$, is a parameter matrix that needs to be learned. To furtherly capture 161 the relative importance of the *i*-th driving variable compared to other driving variables, 162 variable attention is used for the summarized information h_{sum}^i and h_T^i . Note that h_T^i 163 is also a kind of summarized information derived by the LSTM [*Guo et al.*, 2019; 164 *Hochreiter and Schmidhuber*, 1997]. The weight or importance of the *i*-th driving 165 variable w_i is calculated as:

$$w_i' = \tanh \left(W_a[h_{sum}^i, h_T^i] \right)$$

$$w_i = \frac{e^{w_i'}}{\sum_{i=1}^n e^{w_j'}}$$
(4)

Finally, using the weighted sum of all driving variables, the model generates the prediction \hat{Y}_{T+1} :

168

$$o_i = W_o[h_{sum}^i, h_T^i] + b_o$$

$$\hat{Y}_{T+1} = \sum_{i=1}^n o_i w_i$$
(5)

169 where $W_a \in R^{1 \times 2m}$, is a learnable parameter matrix and the linear function with weight 170 $W_o \in R^m$ and bias $b_o \in R$, along with attention calculated weight w_i , produce the 171 final prediction result. The parameters of attention-based LSTM are learned via a back-172 propagation algorithm by minimizing the mean-squared error between predictions and 173 observations [*Guo et al.*, 2019; *Leung and Haykin*, 1991]. 174 The AttentionFire model is implemented with python under Python 3 environment.

175 The model is open-access at https://zenodo.org/record/6903284#.YvH8F-zMJmP under





- 176 Creative Commons Attribution 4.0 International license. Detailed code and descriptions
- 177 are included in the repository including loading datasets, model initialization, training,
- 178 predicting, saving parameters, and loading the trained model (see more details in code
- 179 availability section).
- 180 2.2 Baseline models and model settings

Five other widely used Machine learning (ML) models are used as baseline models 181 182 to compare with AttentionFire model: random forest (RF) [Coffield et al., 2019; Gray 183 et al., 2018b], decision tree (DT) [Amatulli et al., 2006], gradient boosting decision tree 184 (GBDT) [Coffield et al., 2019], artificial neuro network (ANN) [Joshi and Sukumar, 2021; Zhu et al., 2021], and naive LSTM. The inputs of climate and fuel-related 185 variables for the first four models (non-sequence models) are variables of the latest 186 three month available for prediction [Yu et al., 2020] while the corresponding inputs of 187 188 naive LSTM and AttentionFire models are whole-year historical time sequences which cover dynamics from wet to dry seasons to capture short- and long-term dependencies 189 underlying the input sequence[Guo et al., 2019; Li et al., 2020; Qin et al., 2017; 190 191 Vaswani et al., 2017]. The socioeconomic predictors (i.e., population, road density, livestock) consider only the more recent and available statistics typically reported at a 192 year scale. For each model, we iteratively leave one-year dataset out for testing and use 193 the remaining dataset for model training and validation. Details of the settings for used 194 195 models in experiments are listed in Table S1.

196 **2.3 Datasets**

197 Satellite-based global burned area dataset (Global Fire Emissions Database [Giglio et





al., 2013]) is used as prediction target, and datasets of various socio-environmental 198 199 drivers are used as model inputs. Population density, livestock density, road-network 200 density, and land use are considered as anthropogenic factors on fire ignition and spread. Fuel variables include fuel moisture, live and dead vegetation biomass. Seven 201 202 meteorology variables from NCEP-DOE Reanalysis are considered, including air temperature, precipitation, surface pressure, wind speed, specific humidity, downward 203 204 shortwave radiation, and vapor pressure deficit. Details of each dataset and 205 corresponding references are listed in Table S2. The raw datasets were unified to the 206 same spatial resolution (T62 resolution: 94×192) with a covering period from 1997 to 2015. 207

For future projection (2016-2055) of burned area with AttentionFire model, land 208 209 use changes [Hurtt et al., 2020], population growth, projected climate and fuel from 210 fully coupled CESM simulation under high emission scenario (ssp585) were used as the ML model input. The reason to select 2016-2055 as the projected period was that 211 during 2016-2055 99th percentiles of precipitation, temperature, and vapor pressure 212 213 deficit were within the range of corresponding historical observations, which means that the trained model has covered the range of most projected drivers in the near-future 214 and can alleviate extrapolation uncertainty caused by climate change. 215

216

217 3 Results and Discussions

218 **3.1 Model predictability on burned area spatial-temporal dynamics**

219 The AttentionFire model accurately captured the spatial distribution and temporal





220	variations (Fig. 2 and Fig. S1) of wildfire burned areas over NHAF, SHAF, and SHSA
221	regions. The AttentionFire model had the lowest mean absolute errors (MAE) between
222	model predicted and observed grided monthly burned areas among the six ML
223	approaches. The gridded mean absolute errors of burned area for AttentionFire were
224	110, 142, and 39 Kha yr ⁻¹ in NHAF, SHAF, and SHSA regions, which were respectively
225	6%~66%, 13%~65%, and 11%~42% lower than the other 5 ML approaches in the three
226	regions. These results highlight the capability of the AttentionFire model to capture
227	critical driving factors of burned area across time and space.
228	The fact that the AttentionFire model outperformed the other five models (Fig. 2g-
229	i) indicates the benefit of skillfully integrating time-lagged and spatially heterogenous
230	controls from critical drivers on wildfires. Compared to non-sequence models (i.e., RF,
231	MLP, DT, and GBDT), the AttentionFire model adaptively captured historical

232 dependencies of wildfires on climate conditions from wet to dry seasons [*Andela and*

233 Van Der Werf, 2014; Archibald et al., 2009; Chen et al., 2011; Van Der Werf et al., 2008]

(more detailed analysis is provided in next section). Compared to the naive LSTM
models, the variable and temporal attention mechanisms integrated in AttentionFire has
proven to be beneficial to model performance.

The spatial heterogeneity and temporal variation of wildfire responses to complex environmental and human factors have made wildfire predictions challenging, especially at large spatial scales [*Andela and Van Der Werf*, 2014; *Chen et al.*, 2016; *Chen et al.*, 2011; *Littell et al.*, 2016; *Zhou et al.*, 2020]. The capability of the AttentionFire model to reasonably predict spatial and temporal distributions of burned





- 242 area ahead of fire season allows more time to explore and implement management
- 243 options, such as allocation of firefighting resources, fuel clearing or targeted burning
- restrictions [*Chen et al.*, 2011].



Fig. 2. The AttentionFire model accurately captured burned area spatial dynamics. (af) Spatial distribution of observed and AttentionFire predicted fire season mean burned
areas with one-month lead time in Northern Hemisphere Africa (NHAF), Southern
Hemisphere Africa (SHAF), and Southern Hemisphere South America (SHSA) regions.
(g-i) Performance (in terms of mean absolute error) of AttentionFire and other five
baseline models.

252 **3.2 Dominant drivers of tropical burned area dynamics**





253	The AttentionFire model dynamically weights variable importance and highlights
254	critical temporal windows [Guo et al., 2019; Li et al., 2020; Liang et al., 2018; Qin et
255	al., 2017; Vaswani et al., 2017] that maximize model predictability. Therefore, the
256	variable weights could inform dominant physical processes, while the temporal weights
257	reflect the temporal dependency structure, making it interpretable for spatial-temporal
258	analysis. For the AttentionFire model predictions, the variable weights showed that
259	climate wetness exerted strong and spatial heterogenous controls on burned areas.
260	Specifically, precipitation (for SHAF and SHSA regions) and vapor pressure deficit
261	(VPD; for NHAF region) played the most important roles (Fig. 3) in burned area
262	prediction during fire seasons (defined as the four months with the largest burned areas,
263	Fig. S2), and the control strengths from those climate wetness variables on fires were
264	significantly (one-tailed t-test, p-value<0.05) stronger in regions with larger burned
265	areas (gridcells with top 10% burned areas) than those with smaller burned areas
266	(gridcells with last 90% burned areas) (Fig. 4 a-f).











- 274 normalized.
- 275

276	In AttentionFire model predictions, the precipitation and VPD explained $\sim 66\%$ -
277	\sim 80% (Fig. S3) of the annual mean fire season wildfire burned areas. Variations of VPD
278	and precipitation not only affect fire season ignition likelihood and fire spread [Holden
279	et al., 2018; Sedano and Randerson, 2014] through fuel moisture, but also regulate
280	vegetation growth, fuel structure [Gale et al., 2021] (e.g., fuel composition and spatial
281	connectivity), and fuel availability [Littell et al., 2009; Littell et al., 2016; Mueller et
282	al., 2020; Van Der Werf et al., 2008]. The importance of these climate wetness variables
283	confirms the dominant roles of local water balances and air dryness for wildfire
284	prediction from sub-seasonal to seasonal scales [Archibald et al., 2009; Chen et al.,
285	2011; Littell et al., 2016], especially in regions with large burned areas.

286 Furthermore, we found that the emergent functional relationships between climate wetness and wildfire burned area were parabolic (Fig. S3): i.e., enhancement of 287 historical precipitation or decline of historical VPD (indicating wetter conditions) first 288 increased burned area in more xeric conditions, then suppressed burned area under more 289 mesic conditions, consistent with previous findings in subtropical regions [Andela and 290 Van Der Werf, 2014; Van Der Werf et al., 2008]. The transition points of these emergent 291 functional relationships (thresholds at which the relationships reverse) were region 292 specific, and these relationships may be useful for developing, tuning, and 293 benchmarking wildfire models [Zhu et al., 2021]. 294





295	For the time lags between those dominant climate wetness variables and fire-season
296	burned areas, our results demonstrated that burned area over NHAF was more
297	modulated by relatively short-term wetness (VPD during wet-to-dry and onset of dry
298	season, from September to December), while SHAF and SHSA burned areas depended
299	more on long-term wetness (precipitation during wet and wet-to-dry season, December
300	to March in SHAF, and November to April in SHSA) (Fig. 4g-i). The short-term
301	variations of climate wetness can directly affect near-surface temperature and moisture
302	availability, which affect fuel flammability [Holden et al., 2018; Littell et al., 2016],
303	while the long-term wetness (e.g., during rainy season) can affect fuel availability,
304	composition, and spatial connectivity, which can result in even stronger long time-
305	lagged controls on dry-season burned areas [Abatzoglou and Kolden, 2013; Andela and
306	Van Der Werf, 2014; Archibald et al., 2009; Chen et al., 2011; Littell et al., 2016; Van
307	<i>Der Werf et al.</i> , 2008].
200	

308

309







Fig. 4. Spatial-temporal importance of climate wetness variables for burned area dynamics. (a-c) Spatial importance of climate wetness variables for fire-season burned areas. (d-f) statistical comparison of the climate wetness variable importance over regions with large and small burned areas. (g-i) fire season burned area dependency on the history of the climate wetness driver over Northern Hemisphere Africa (NHAF), Southern Hemisphere Africa (SHAF), and Southern Hemisphere South America (SHSA) regions.

Previous work has shown that when and where fires occurred during dry season can be affected by precipitation induced fuel availability patterns during wet and during wet-to-dry transition seasons in savannah ecosystems [*Andela and Van Der Werf*, 2014; *Archibald et al.*, 2009; *Van Der Werf et al.*, 2008]. Also, precipitation variations during wet and wet-to-dry transition seasons in the tropical forest ecosystem can affect soil recharge during wet season and further affect plant transpiration, local surface humidity,





323	and precipitation during the following dry season [Chen et al., 2011; Malhi et al., 2008;
324	Ramos da Silva et al., 2008]. The exact responses of fires to short-and-long term climate
325	variations depend on both local wetness and fuel conditions (e.g., fires in wetter
326	ecosystems with enough fuel availability can be mainly limited by the length of dry
327	season, while fires in drier ecosystems can be limited by fuel availability during wet
328	season [Andela and Van Der Werf, 2014; Van Der Werf et al., 2008]). Therefore, an
329	effective way of integrating the climate wetness history (i.e., AttentionFire model) can
330	lead to more accurate predictions of burned area spatial-temporal dynamics.
331	3.3 Possible usage of oceanic index for long-leading time predictions
332	In ASA regions, large-scale variations of oceanic dynamics can directly influence
332 333	In ASA regions, large-scale variations of oceanic dynamics can directly influence local climate (e.g., precipitation variations during wet seasons [<i>Andela and Van Der</i>
333	local climate (e.g., precipitation variations during wet seasons [Andela and Van Der
333 334	local climate (e.g., precipitation variations during wet seasons [Andela and Van Der Werf, 2014; Chen et al., 2011]) through time-lagged controls of teleconnections and
333 334 335	local climate (e.g., precipitation variations during wet seasons [<i>Andela and Van Der Werf</i> , 2014; <i>Chen et al.</i> , 2011]) through time-lagged controls of teleconnections and indirectly influence fires during following dry seasons [<i>Andela et al.</i> , 2017; <i>Chen et al.</i> ,
333 334 335 336	local climate (e.g., precipitation variations during wet seasons [<i>Andela and Van Der</i> <i>Werf</i> , 2014; <i>Chen et al.</i> , 2011]) through time-lagged controls of teleconnections and indirectly influence fires during following dry seasons [<i>Andela et al.</i> , 2017; <i>Chen et al.</i> , 2016; <i>Chen et al.</i> , 2011]. Therefore, we hypothesized that ocean dynamics might benefit
333 334 335 336 337	local climate (e.g., precipitation variations during wet seasons [<i>Andela and Van Der</i> <i>Werf</i> , 2014; <i>Chen et al.</i> , 2011]) through time-lagged controls of teleconnections and indirectly influence fires during following dry seasons [<i>Andela et al.</i> , 2017; <i>Chen et al.</i> , 2016; <i>Chen et al.</i> , 2011]. Therefore, we hypothesized that ocean dynamics might benefit AttentionFire model predictions, especially for long leading time fire predictions

We compared model performance for short term (1-4 month ahead), and long term 341 (5-8 month ahead) fire predictions with and without considering the four oceanic 342 indexes. Relative to the MAE of short-term predictions, the mean MAE of long-term 343 predictions without and with teleconnections increased by ~34% and ~14% in NHAF, 344

352





345 ~34% and ~15% in SHAF, and ~17% and ~7% in SHSA, respectively, indicating the 346 decline of system predictability with longer leading time (Figure 5). However, for long-347 term predictions, including oceanic indexes and teleconnections could decrease the 348 mean MAE by ~20%, ~19%, and ~11% in NHAF, SHAF, and SHSA regions, 349 respectively, compared with the case without oceanic indexes. The results demonstrated 350 the potential usage of teleconnections for longer than 5 months leading time burned 351 area predictions.



Fig. 5: Performance of AttentionFire burned area predictions with 1~4 month leading
time (short-term) and with 5-8 month leading (long-term). MAE is mean absolute error.
Long term prediction with OI means the AttentionFire model also considered four
ocean indices that have been widely used for fire prediction over South American and
African regions. Four OIs are: Oceanic Niño Index (ONI), Atlantic multidecadal
Oscillation (AMO) index, Tropical Northern Atlantic (TNA) Index, and Tropical
Southern Atlantic (TSA) Index.

360 3.4 Future trends of burned area over Africa and South America

Due to climate change and human activities [*Andela et al.*, 2017], strong but opposing trends of burned areas have been observed in Northern (decreasing) and Southern (increasing) Hemisphere Africa [*Andela and Van Der Werf*, 2014], and within different regions of Southern Hemisphere America [*Andela et al.*, 2017] during the





365	recent two decades, resulting in an overall declining burned area trend in Africa and
366	South America. However, whether this decline will persist is under debate. On one hand,
367	the projected increases in population, expansion of agriculture, mechanized (fire-free)
368	management, and fire suppression policies will likely continue to decrease burned areas
369	[Andela and Van Der Werf, 2014] (e.g., human activities were regarded as one of the
370	main drivers for fire decline in NHAF region). On the other hand, future climate change
371	[Dai, 2013; Taufik et al., 2017] could outweigh human impacts and result in
372	unprecedented fire-prone environments in the tropics [Malhi et al., 2008; Pechony and
373	Shindell, 2010] (e.g., fires showed strong dependency with climate wetness in NHAF,
374	SHAF [Andela and Van Der Werf, 2014; Archibald et al., 2009] and SHSA [Chen et al.,
375	2011] regions).

376

377







378 Fig. 6: Future burned area trends under the SSP585 high emission scenario. (a-c) spatial 379 distribution of fire season burned area trends using drivers with interannual variations; dots in (a-c) indicate gridcells with statistically significant changes in the trend. (d-f) 380 regionally aggregated burned area changes with historical mean subtracted. Blue and 381 382 red lines respectively represent burned area anomaly in history and future; the black line represents future burned area trend while removing the interannual variations of 383 the dominant variable. Solid lines represented significant BA trends (p value <0.05) 384 while dashed lines represented non-significant BA trends. 385

Considering land use changes, population growth, and projected climate under the SSP585 high emission scenario, our model predicted that burned areas in the NHAF region will continue to decline; the currently increasing trend will be dampened in the





SHAF region, and the currently decreasing trend will be reversed in SHSA region (Fig.
6). Over NHAF and SHSA, burned area trends at the gridcell level are mostly robust
(Fig. 6a-c; *p* value<0.05) and of the same sign, thus resulting in a robust trend at
regional scale.

393 To investigate what drives future burned area changes, we iteratively surrogated each driver with its climatology while keeping the other factors the same. Gridded 394 395 burned area changing trends in NHAF and SHSA were mostly affected by VPD changes 396 (Table S3), and removing VPD inter-annual changes resulted in non-significant burned 397 area trends at the whole NHAF and SHSA region (Fig. 6). VPD was projected to continuously increase due to warming but had different implications over NHAF and 398 SHSA. Over the relatively fuel abundant SHSA region, increased VPD will likely 399 increase burned area (Pearson r= 0.45, p value <0.05, Fig. S4) through increasing fuel 400 401 dryness and combustibility [Chen et al., 2011; Kelley et al., 2019; Malhi et al., 2008; Van Der Werf et al., 2008]. In contrast, over the semi-arid savannah dominated NHAF 402 403 region (less fuel, compared with SHSA), higher VPD could decrease burned area (Pearson r= -0.81, p value <0.05, Fig. S4) through limiting plant growth and fuel 404 availability [Andela et al., 2017; Andela and Van Der Werf, 2014; Van Der Werf et al., 405 2008]. For the SHAF, population growth followed by climate changes (Table S3) 406 showed stronger influences on grided burned area changes while the heterogeneity of 407 408 wildfire responses finally led to a non-significant trend at the regional scale (Fig. 6). Our findings highlight the importance of climate changes on understanding future 409 burned area dynamics, and motivate better representation of climate wetness effects on 410





- 411 wildfire dynamics in process-based and machine learning-based wildfire prediction
- 412 models.
- 413
- 414 4. Conclusions

415 This study developed an interpretable machine learning model (AttentionFire v1.0) for burned areas predictions over African and south American regions. Compared with 416 417 observations and other five widely used machine learning baseline models, we 418 demonstrated the effectiveness of the AttentionFire model to capture the magnitude, 419 spatial distribution, and temporal variation of burned areas. "Attention" mechanisms enabled the interpretation of complex but critical spatial-temporal patterns [Guo et al., 420 2019; Li et al., 2020; Liang et al., 2018; Qin et al., 2017; Vaswani et al., 2017], thus 421 422 uncovering the black-boxed relationships in machine learning models for burned area 423 predictions. We demonstrated the spatiotemporally heterogenous and strong timelagged controls from local climate wetness on burned areas. Furthermore, under the 424 SSP585 high emission scenario, our results suggested that the increasing trend in 425 426 burned area over southern Africa will dampened, and the declining trend in burned area over fuel abundant southern America will reverse. This study highlights the importance 427 of skillful representation of spatiotemporally heterogenous and strong time-lagged 428 climate wetness effects on understanding wildfire dynamics and developing advanced 429 430 early fire warning models.

431

432 Acknowledgements:





433	This research was supported by the Director, Office of Science, Office of Biological
434	and Environmental Research of the US Department of Energy under contract no.
435	DEAC02-05CH11231 as part of their Regional and Global Climate Modeling program
436	through the Reducing Uncertainties in Biogeochemical Interactions through Synthesis
437	and Computation Scientific Focus Area (RUBISCO SFA) project and as part of the
438	Energy Exascale Earth System Model (E3SM) project.

439

440 Code availability

The source code of AttentionFire_v1.0 is archive at Zenodo repository: 441 https://zenodo.org/record/6903284#.YuA3oOzMIaF, under Creative 442 Commons Attribution 4.0 International license, with four zip files: data, data preparation, model, 443 and example. The "data" file contains the links to all raw datasets used to drive the 444 445 model (e.g., burned areas, climate forcing). The "data preparation" file contains the code to preprocess the raw datasets and make them be ready for training and testing the 446 AttentionFire model. The "model" file contains the python code of AttentionFire model. 447 448 The "example" file gives a detailed example of how to use the AttentionFire model for burned area predictions. 449

There is also a tutorial file "Data_Model_Tutorial" that contain descriptions on (1) how to load the raw datasets; (2) how to prepare the input and output datasets for ML model; (3) how to initialize the ML model and run the model (4) how to train the ML model and use the trained ML model for predictions; (5) how to save and load the model parameters and save the predicted results.





455

- 456 Data availability
- 457 Burned area: Global Fire Emissions Database
- 458 https://daac.ornl.gov/VEGETATION/guides/fire_emissions_v4.html
- 459 NCEP-DOE Reanalysis Climate forcings:
- 460 https://psl.noaa.gov/data/gridded/data.ncep.reanalysis2.html
- 461 **Population:** https://landscan.ornl.gov/
- 462 Road density: https://www.globio.info/download-grip-dataset
- 463 Livestock density: https://www.fao.org/dad-is/en/
- 464 Land cover change: https://luh.umd.edu/data.shtml
- 465 Oceanic index: https://psl.noaa.gov/data/climateindices/list/
- 466

467 Author contributions

- 468 QZ and FL designed the study. QZ, FL, and MC designed the model experiments. FL
- 469 wrote the code and ran the experiments. LZ, WR, JR, LX, HW, ZG, and JG all
- 470 contributed to the interpretation of the results and writing of the paper.
- 471

472 References

Abatzoglou, J. T., and C. A. Kolden (2013), Relationships between climate and macroscale area burned 473 474 in the western United States, International Journal of Wildland Fire, 22(7), 1003-1020. 475 Amatulli, G., M. J. Rodrigues, M. Trombetti, and R. Lovreglio (2006), Assessing long-term fire risk at 476 local scale by means of decision tree technique, Journal of Geophysical Research: Biogeosciences, 477 111(G4). 478 Andela, N., D. C. Morton, L. Giglio, Y. Chen, G. Van Der Werf, P. S. Kasibhatla, R. DeFries, G. Collatz, 479 S. Hantson, and S. Kloster (2017), A human-driven decline in global burned area, Science, 480 356(6345), 1356-1362.





481	Andela, N., and G. R. Van Der Werf (2014), Recent trends in African fires driven by cropland expansion
482	and El Niño to La Niña transition, <i>Nature Climate Change</i> , 4(9), 791-795.
483	Aragao, L. E. O., Y. Malhi, N. Barbier, A. Lima, Y. Shimabukuro, L. Anderson, and S. Saatchi (2008),
484	Interactions between rainfall, deforestation and fires during recent years in the Brazilian Amazonia,
485	Philosophical Transactions of the Royal Society B: Biological Sciences, 363(1498), 1779-1785.
486	Archibald, S., D. P. Roy, B. W. van Wilgen, and R. J. Scholes (2009), What limits fire? An examination
487	of drivers of burnt area in Southern Africa, Global Change Biology, 15(3), 613-630.
488	Benavides-Solorio, J., and L. H. J. H. P. MacDonald (2001), Post-fire runoff and erosion from simulated
489	rainfall on small plots, Colorado Front Range, 15(15), 2931-2952.
490	Bowman, D. M., J. K. Balch, P. Artaxo, W. J. Bond, J. M. Carlson, M. A. Cochrane, C. M. D'Antonio,
491	R. S. DeFries, J. C. Doyle, and S. P. Harrison (2009), Fire in the Earth system, science, 324(5926),
492	481-484.
493	Chen, Y., D. C. Morton, N. Andela, L. Giglio, and J. T. Randerson (2016), How much global burned area
494	can be forecast on seasonal time scales using sea surface temperatures?, Environmental Research
495	Letters, 11(4), 045001.
496	Chen, Y., J. T. Randerson, S. R. Coffield, E. Foufoula-Georgiou, P. Smyth, C. A. Graff, D. C. Morton, N.
497	Andela, G. R. van der Werf, and L. Giglio (2020), Forecasting global fire emissions on subseasonal
498	to seasonal (S2S) time scales, Journal of advances in modeling earth systems, 12(9),
499	e2019MS001955.
500	Chen, Y., J. T. Randerson, D. C. Morton, R. S. DeFries, G. J. Collatz, P. S. Kasibhatla, L. Giglio, Y. Jin,
501	and M. E. Marlier (2011), Forecasting fire season severity in South America using sea surface
502	temperature anomalies, Science, 334(6057), 787-791.
503	Coffield, S. R., C. A. Graff, Y. Chen, P. Smyth, E. Foufoula-Georgiou, and J. T. Randerson (2019),
504	Machine learning to predict final fire size at the time of ignition, International journal of wildland
505	fire.
506	Dai, A. (2013), Increasing drought under global warming in observations and models, <i>Nature climate</i>
507	<i>change</i> , <i>3</i> (1), 52-58.
508	Etminan, M., G. Myhre, E. Highwood, and K. J. G. R. L. Shine (2016), Radiative forcing of carbon
509	dioxide, methane, and nitrous oxide: A significant revision of the methane radiative forcing, 43(24),
510	12,614-612,623.
511	Gale, M. G., G. J. Cary, A. I. Van Dijk, and M. Yebra (2021), Forest fire fuel through the lens of remote
512	sensing: Review of approaches, challenges and future directions in the remote sensing of biotic
513	determinants of fire behaviour, Remote Sensing of Environment, 255, 112282.
514	Giglio, L., J. T. Randerson, and G. R. Van Der Werf (2013), Analysis of daily, monthly, and annual burned
515	area using the fourth-generation global fire emissions database (GFED4), Journal of Geophysical
516	Research: Biogeosciences, 118(1), 317-328.
517	Gray, M. E., L. J. Zachmann, and B. G. Dickson (2018a), A weekly, continually updated dataset of the
518	probability of large wildfires across western US forests and woodlands, Earth System Science Data,
519	10(3), 1715-1727.
520	Gray, M. E., L. J. Zachmann, and B. G. J. E. S. S. D. Dickson (2018b), A weekly, continually updated
521	dataset of the probability of large wildfires across western US forests and woodlands, 10(3), 1715-
522	1727.
523	Guo, T., T. Lin, and N. Antulov-Fantulin (2019), Exploring interpretable LSTM neural networks over
524	multi-variable data, paper presented at International Conference on Machine Learning, PMLR.





525 526	Hantson, S., A. Arneth, S. P. Harrison, D. I. Kelley, I. C. Prentice, S. S. Rabin, S. Archibald, F. Mouillot,S. R. Arnold, and P. Artaxo (2016), The status and challenge of global fire modelling,
527	Biogeosciences, 13(11), 3359-3375.
528	Hochreiter, S., and J. Schmidhuber (1997), Long short-term memory, <i>Neural computation</i> , 9(8), 1735-
520 529	1780.
530	Holden, Z. A., A. Swanson, C. H. Luce, W. M. Jolly, M. Maneta, J. W. Oyler, D. A. Warren, R. Parsons,
530 531	and D. Affleck (2018), Decreasing fire season precipitation increased recent western US forest
532	wildfire activity, <i>Proceedings of the National Academy of Sciences</i> , 115(36), E8349-E8357.
533	Hurtt, G. C., L. Chini, R. Sahajpal, S. Frolking, B. L. Bodirsky, K. Calvin, J. C. Doelman, J. Fisk, S.
534	Fujimori, and K. Klein Goldewijk (2020), Harmonization of global land use change and
535	management for the period 850–2100 (LUH2) for CMIP6, Geoscientific Model Development,
536	<i>13</i> (11), 5425-5464.
537	Jain, P., S. C. Coogan, S. G. Subramanian, M. Crowley, S. Taylor, and M. D. Flannigan (2020), A review
538	of machine learning applications in wildfire science and management, <i>Environmental Reviews</i> ,
539	<i>28</i> (4), 478-505.
540	Joshi, J., and R. Sukumar (2021), Improving prediction and assessment of global fires using multilayer
541	neural networks, <i>Scientific reports</i> , <i>11</i> (1), 1-14.
542	Kelley, D. I., I. Bistinas, R. Whitley, C. Burton, T. R. Marthews, and N. Dong (2019), How contemporary
543	bioclimatic and human controls change global fire regimes, <i>Nature Climate Change</i> , 9(9), 690-696.
544	Knorr, W., F. Dentener, JF. Lamarque, L. Jiang, and A. Arneth (2017), Wildfire air pollution hazard
545	during the 21st century, Atmospheric Chemistry and Physics, 17(14), 9223-9236.
546	Lelieveld, J., J. S. Evans, M. Fnais, D. Giannadaki, and A. Pozzer (2015), The contribution of outdoor
547	air pollution sources to premature mortality on a global scale, Nature, 525(7569), 367-371.
548	Leung, H., and S. Haykin (1991), The complex backpropagation algorithm, <i>IEEE Transactions on signal</i>
549	processing, 39(9), 2101-2104.
550	Li, F., Z. Gui, Z. Zhang, D. Peng, S. Tian, K. Yuan, Y. Sun, H. Wu, J. Gong, and Y. Lei (2020), A
551	hierarchical temporal attention-based LSTM encoder-decoder model for individual mobility
552	prediction, Neurocomputing, 403, 153-166.
553	Liang, Y., S. Ke, J. Zhang, X. Yi, and Y. Zheng (2018), Geoman: Multi-level attention networks for geo-
554	sensory time series prediction, paper presented at IJCAI.
555	Littell, J. S., D. McKenzie, D. L. Peterson, and A. L. Westerling (2009), Climate and wildfire area burned
556	in western US ecoprovinces, 1916–2003, Ecological Applications, 19(4), 1003-1021.
557	Littell, J. S., D. L. Peterson, K. L. Riley, Y. Liu, and C. H. Luce (2016), A review of the relationships
558	between drought and forest fire in the United States, Global change biology, 22(7), 2353-2369.
559	Malhi, Y., J. T. Roberts, R. A. Betts, T. J. Killeen, W. Li, and C. A. Nobre (2008), Climate change,
560	deforestation, and the fate of the Amazon, science, 319(5860), 169-172.
561	Mueller, S. E., A. E. Thode, E. Q. Margolis, L. L. Yocom, J. D. Young, and J. M. Iniguez (2020), Climate
562	relationships with increasing wildfire in the southwestern US from 1984 to 2015, Forest Ecology
563	and Management, 460, 117861.
564	Pechony, O., and D. T. Shindell (2010), Driving forces of global wildfires over the past millennium and
565	the forthcoming century, Proceedings of the National Academy of Sciences, 107(45), 19167-19170.
566	Qin, Y., D. Song, H. Chen, W. Cheng, G. Jiang, and G. Cottrell (2017), A dual-stage attention-based
567	recurrent neural network for time series prediction, arXiv preprint arXiv:1704.02971.
568	Rabin, S. S., J. R. Melton, G. Lasslop, D. Bachelet, M. Forrest, S. Hantson, J. O. Kaplan, F. Li, S.





569	Mangeon, and D. S. Ward (2017), The Fire Modeling Intercomparison Project (FireMIP), phase 1:
570	experimental and analytical protocols with detailed model descriptions, Geoscientific Model
571	Development, 10(3), 1175-1197.
572	Ramanathan, V., P. Crutzen, J. Kiehl, and D. Rosenfeld (2001), Aerosols, climate, and the hydrological
573	cycle, science, 294(5549), 2119-2124.
574	Ramos da Silva, R., D. Werth, and R. Avissar (2008), Regional impacts of future land-cover changes on
575	the Amazon basin wet-season climate, Journal of climate, 21(6), 1153-1170.
576	Randerson, J. T., H. Liu, M. G. Flanner, S. D. Chambers, Y. Jin, P. G. Hess, G. Pfister, M. Mack, K.
577	Treseder, and L. J. s. Welp (2006), The impact of boreal forest fire on climate warming, 314(5802),
578	1130-1132.
579	Reichstein, M., G. Camps-Valls, B. Stevens, M. Jung, J. Denzler, and N. Carvalhais (2019), Deep
580	learning and process understanding for data-driven Earth system science, Nature, 566(7743), 195-
581	204.
582	Sedano, F., and J. Randerson (2014), Multi-scale influence of vapor pressure deficit on fire ignition and
583	spread in boreal forest ecosystems, Biogeosciences, 11(14), 3739-3755.
584	Shvetsov, E. G., E. A. Kukavskaya, L. V. Buryak, and K. J. E. R. L. Barrett (2019), Assessment of post-
585	fire vegetation recovery in Southern Siberia using remote sensing observations, 14(5), 055001.
586	Taufik, M., P. J. Torfs, R. Uijlenhoet, P. D. Jones, D. Murdiyarso, and H. A. Van Lanen (2017),
587	Amplification of wildfire area burnt by hydrological drought in the humid tropics, Nature Climate
588	Change, 7(6), 428-431.
589	Teckentrup, L., S. P. Harrison, S. Hantson, A. Heil, J. R. Melton, M. Forrest, F. Li, C. Yue, A. Arneth,
590	and T. Hickler (2019), Response of simulated burned area to historical changes in environmental
591	and anthropogenic factors: a comparison of seven fire models, <i>Biogeosciences</i> , 16(19), 3883-3910.
592	Turco, M., S. Jerez, F. J. Doblas-Reyes, A. AghaKouchak, M. C. Llasat, and A. Provenzale (2018), Skilful
593	forecasting of global fire activity using seasonal climate predictions, Nature communications, 9(1),
594	1-9.
595	Van Der Werf, G. R., J. T. Randerson, L. Giglio, N. Gobron, and A. Dolman (2008), Climate controls on
596	the variability of fires in the tropics and subtropics, <i>Global Biogeochemical Cycles</i> , 22(3).
597	Vaswani, A., N. Shazeer, N. Parmar, J. Uszkoreit, L. Jones, A. N. Gomez, L. Kaiser, and I. Polosukhin
598	(2017), Attention is all you need, arXiv preprint arXiv:1706.03762.
599	Werf, G. R., J. T. Randerson, L. Giglio, T. T. v. Leeuwen, Y. Chen, B. M. Rogers, M. Mu, M. J. Van
600	Marle, D. C. Morton, and G. J. J. E. S. S. D. Collatz (2017), Global fire emissions estimates during
601	1997–2016, 9(2), 697-720.
602	Xu, X., G. Jia, X. Zhang, W. J. Riley, and Y. Xue (2020), Climate regime shift and forest loss amplify
603	fire in Amazonian forests, Global Change Biology, 26(10), 5874-5885.
604	Yu, Y., J. Mao, P. E. Thornton, M. Notaro, S. D. Wullschleger, X. Shi, F. M. Hoffman, and Y. Wang
605	(2020), Quantifying the drivers and predictability of seasonal changes in African fire, Nature
606	communications, 11(1), 1-8.
607	Zhou, W., D. Yang, SP. Xie, and J. J. N. C. C. Ma (2020), Amplified Madden–Julian oscillation impacts
608	in the Pacific–North America region, Nature Climate Change, 10(7), 654-660.
609	Zhu, Q., F. Li, W. J. Riley, L. Xu, L. Zhao, K. Yuan, H. Wu, J. Gong, and J. T. Randerson (2021), Building
610	a machine learning surrogate model for wildfire activities within a global earth system model,
611	Geoscientific Model Development Discussions, 1-22.
612	