ForamEcoGENIE 2.0: Incorporating symbiosis and spine traits into a trait-based global planktic foraminifera model

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Abstract. Planktic foraminifera are major marine calcifiers in the modern ocean regulating the marine inorganic carbon pump and generating marine fossil archives of past climate change. Some planktic foraminifera evolved spine and symbiosis, increasing functional trait diversity and expanding their ecological niches. Here we incorporate symbiosis and spine traits into

- 10 the ForamEcoGENIE model, a trait-based model focusing on functional traits rather than individual species, to enable us to study the importance of foraminifera biodiversity in the palaeoceanographic environment. We calibrated the modelled new traits using Latin Hypercube Sampling. We identified the <u>optimal model parameters</u> from an ensemble of 1200 runs compared with observations from global core-tops, sediment traps, and plankton nets. The model successfully captures the global distribution and seasonal variation of the 4 major functional groups including dominance of the symbiont-obligate type in
- 15 subtropical gyres and the symbiont-barren type in the productive subpolar oceans. The global annual mean biomass and foraminifer-derived carbon export rate is correctly predicted, with biomass from 0.001 to 0.010 mmol C m⁻³ and organic carbon export 0.002-0.031 mmol C m⁻² d⁻¹, The seasonal peak time of biomass and organic carbon export are also generally captured, although the amplitude is not able to be reproduced in productive area. Both the sparseness of observations and the model's
- limitation in <u>upwelling regions</u> likely contributes to the discrepancy. Our model overcomes the lack of biodiversity in previous
 version and offers the potential to explore foraminifera ecology dynamics and its impact on biogeochemistry in modern, future and paleogeographic environments.

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Deleted: Our model approximates a $3.05 \text{ g m}^2 \text{ yr}^1$ global mean foraminifer-derived calcite flux and 1.1 Gt yr^1 total calcite export, account for 19% of the global pelagic marine calcite budget within the lower range of modern calcite estimates. The calcite export is mostly derived from the symbiont-barren non-spinose group (39%) and the symbiont-obligate spinose group (13%).

1 Introduction

Planktic foraminifera are marine calcifying zooplankton that have populated the surface ocean since the mid-Jurassic (~175 Ma). They produce calcite shells (or "tests"), preserved in vast amounts in sediments, that provide proxy archives (e.g., ¹³C, ¹⁸O, Mg/Ca) to reconstruct past climate conditions (Tierney et al., 2020), ocean carbonate chemistry (Hönisch et al., 2012) and

45 study biotic response to environmental change (Todd et al., 2020). In the modern oceans, foraminifera contribute 23-56% of the total open-ocean CaCO₃ export (Schiebel, 2002) alongside the other major calcifiers, coccolithophores (Daniels et al., 2018) and pteropods (Buitenhuis et al., 2019). However, understanding the impacts of environmental change on foraminifers and their role in the carbon cycle is challenged by their low standing stocks in the surface ocean, a (semi)lunar reproductive cycle driving abundances and difficulties in culturing to ground truth physiology (Schiebel and Hemleben, 2017). Modelling

50 planktic foraminifera and their ecology, therefore, plays a critical role in increasing and testing our understanding of their biological and ecological influence on the marine inorganic carbon cycle and their role as a paleo-proxy carrier,

The recent decades have seen significant developments of global foraminifera models due to the increasing data compilations of foraminifer flux and community structure (Siccha and Kucera, 2017; Buitenhuis et al., 2013; Sunagawa et al., 2020). But

- 55 they are either empirical based or limited to the extant species. For example, Waterson et al., (2017) built a Maxent species distribution models based on sediment core data to study the niche variability from Last Glacial Maximum (LGM) to Holocene. Žarić et al. (2006) constructed the statistical model that correlated hydrographical factors and sediment trap abundance of 18 main species. Such correlative models have limitation on predicting future projection because the assumption of constant correlation with environmental factors can be violated (Buckley et al., 2010). In addition, niche models also are ignorant in
- 60 biological interactions which have an important role in shaping species distribution. On the other hand, although Fraile et al. (2008) and Lombard et al. (2011) have built ecophysiology-based mechanistic models PLAFOM and FORAMCLIM which, successfully reconstructed planktic foraminifer's geographical distribution, seasonal and vertical population dynamics (Kretschmer et al., 2018) and held the abilities to simulate in different climatology like the LGM (Fraile et al., 2009) or projected high-emission future (Roy et al., 2015), these two models are species-based and therefore cannot be applied in the

65 deeper times older than the Miocene with many extinct species and cryptic taxa with novel ecologies (Renaud and Schmidt, 2003). Thus, to fill the model gap of not able to leverage the abundant foraminifer fossil information, a mechanistic model but not limited to any species is needed.

Trait-based plankton models as an alternative to species-based models come into vision under such circumstance. This approach which focuses on organismal traits including morphological and physiological properties instead of taxonomic identities provides a mechanistic way to mimic the complex real ocean ecology by defining traits' functional benefits and costs (i.e., trade-offs) (Zakharova et al., 2019; Kiørboe et al., 2018). Models adopting trait-based framework have successfully resembled the diverse marine community such as cyanobacteria (Follows et al., 2007) and diazotroph (Monteiro et al., 2010).

Deleted: first global-scale prediction on foraminifer flux using a Deleted: technique Deleted: Deleted: an ecophysiology based dynamic Deleted: based on CESM (Hurrell et al., 2013) and PISCES (Aumont et al., 2015), respectively. These two models Deleted: production Deleted: in Last Glacial Maximum (LGM) (Fraile et al., 2009), vertical distribution dynamics Deleted: potential diversity change in response to Deleted: projected Deleted: scenarios Deleted: Deleted: However Deleted: , the species-based feature of these models causes limitations for applications. Deleted: For example, parameterisation limits the inclusion of additional species, which are less well understood ecologically. Moreover, while these models are applied to LGM and future oceans, the application to deeper time is severely limited by the existence of extinct species and cryptic taxa with novel ecologies (Renaud and Schmidt, 2003) and non-analogue situations, i.e., time intervals older

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than the Miocene with many none extant species, cannot be assessed

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This strategy is also well suitable for planktic foraminifer because functional traits like body size (Schmidt et al., 2004), size normalised weight (Todd et al., 2020), and symbiosis (Spero and Parker, 1985) have been widely measured and studied. And such functional traits have been consistently evolved in foraminifera across Cenozoic (Ezard et al., 2011). Next we provide the brief introduction for spine and symbiosis traits that determine species-level ecological and functional distinguishment and

115 also compose of the focus of this model study,

<u>Nearly 19 out of 50</u> modern foraminifer species bear eukaryotic algae such as dinoflagellates, chrysophytes and haptophyte as symbionts (Takagi et al., 2019). <u>Compared to non-symbiont species (termed as "symbiont-barren") that live in temperate and polar oceans, symbiotic species dwell in the tropical to subtropical regions (see biogeography in Figure 3) where their</u>

- 120 symbionts are facilitated to photosynthesize and providing an extra energy pathway to foraminifera in the nutrient-deficient environment (LeKieffre et al., 2018; Ortiz et al., 1995; Uhle et al., 1999). The symbiotic group can be further partitioned because some foraminifers that cannot live without their symbionts (termed as "symbiont-obligate") (Bé et al., 1982) while others are flexible (termed as "symbiont-facultative"). As the symbionts affecting latitudinal distribution of foraminifer, the calcareous spine trait shows importance in shaping feeding behaviour. Non-spinose foraminifera rely on rhizopodia to capture
- 125 prey, but spinose foraminifer with spines extruding from the test have larger effective reach range and stronger ability to active prey, which leads to more prey types and sizes like large copepods (Anderson et al., 1979). Laboratory observations show the spinose foraminifera's carnivory preference with the zooplankton/phytoplankton prey protein ratio from 117:1 to 278:1 (Schiebel and Hemleben, 2017, Section 4.1). And the effective encounter rate of a spinose genus can be three orders of magnitude higher than non-spinose foraminifer (Gaskell et al., 2019). Furthermore, roughly half of modern species are spinose,
- 130 therefore this trait is non-negligible. Based on the presence of these two functional traits, the foraminifera can be divided into four functional types: (1) symbiont-barren non-spinose; (2) symbiont-barren spinose; (3) symbiont-facultative non-spinose; (4) and symbiont-obligate spinose (Table 1),

Recently, Grigoratou et al., (2019) developed the first mechanistic and trait-based 0D model for symbiont-barren non-spinose

- 135 foraminifer and coupled it to cGENIE (ForamEcoGENIE <u>1 hereafter</u>, Grigoratou et al., 2021a), an <u>3D</u> Earth System Model of intermediate complexity allowing for fast computational time and widely applied to past climates: Palaeocene–Eocene Thermal Maximum (Ridgwell and Schmidt, 2010), Last Glacial Maximum (Rae et al., 2020) and Cretaceous-Paleogene boundary (Henehan et al., 2019). Such computational efficiency and abundant applications make ForamEcoGENIE easily applicable to a wide range of geological periods with direct links with seawater carbon chemistry and isotope tracers. In this study, we extend ForamEcoGENIE to resolve 3 more critical functional groups of planktic foraminifera by adding the traits of
- symbiosis and spines (tested in Grigoratou et al., 2021b). Thereby, we build a model that can explore foraminifer <u>diversity in</u> past climates (Ezard et al., 2011). We compare the model with three global observational data compilations (core-top, plankton net tow and sediment traps) and test its ability to produce surface biomass, organic carbon and calcite <u>flux</u>, and relative abundance distribution.

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Understanding of foraminifer traits and their functions is crucial for developing a trait-based model. The foremost trait of foraminifer is calcification, with foraminifera building a shell by adding calcite chambers during their development (Caromel et al., 2015). In addition, 19 out of ~50

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Deleted: prefer the low latitude to the symbiont-facultative group which can flexibly bear symbionts and have a wider geography (Figure 1). Another important trait is

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Deleted: (Anderson et al., 1979) These spinose species are mostly symbiotic (except symbiont-barren *Globigerina bulloides* and *Hastigerina pelagica*) and show a preference for omnivorous feeding (Schiebel and Hemleben, 2017). Therefore, traits affect biogeography and trophic activities and lay the foundation of building a trait-based model.⁴

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2 cGENIE ocean and atmosphere physics,

ForamEcoGENIE uses cGENIE (carbon-centric Grid-ENabled Integrated Earth system model), a modular Earth system model of intermediate complexity (EMIC) as the physical ocean simulation framework. The fast climate and ocean physics of

- 185 cGENIE are based on a coarse-resolution 3D frictional geostrophic ocean model coupled to a 2D energy-moisture-balance atmosphere model and a dynamic-thermodynamic sea-ice model (Edwards and Marsh, 2005; Marsh et al., 2011). The ocean has a 36x36 equal-area horizontal grid (uniform in longitude and sine-uniform in latitude) with 16 logarithmically spaced vertical levels as defined in Cao et al., (2009). The physical model is coupled with a model of ocean biogeochemical cycles (Ridgwell et al., 2007; van de Velde et al., 2021), sea-floor sedimentary processes (Ridgwell and Hargreaves, 2007) and marine
- 190 ecosystem processes (Ward et al., 2018). The plankton ecosystem is resolved in the surface layer (0-80.8 m) to mimic light limitation. The model presented in this study is configured with a seasonally forced pre-industrial climate state and a fixed atmospheric CO₂ concentration restored to 278 ppm.

3 Size-based plankton ecosystem framework EcoGENIE

195 <u>3,1, Biogeochemical tracers</u>

The model has three biogeochemical tracers (noted in i_r): carbon, phosphorus, iron. The plankton biomass has same tracers (i.e., C, P, Fe) with notation i_{fo} for each plankton group j_n And phytoplankton or symbiotic foraminifera have an extra state variable for chlorophyll (Chl hereafter). These elements are marked in different colors in Figure 1.

200 3,2 Plankton cell size and quota

In EcoGENIE, individual body size determines key physiological processes, including nutrient uptake, photosynthesis, grazing gain and predation through allometric scaling (West et al., 1997), because of its role as a master trait among pelagic organisms (Andersen et al., 2016). The modelled size-dependent parameters (except for photosynthesis) follow a generic power law: $P = aV^b$ with P the size-based parameter, V the spheric biovolume, and a, and b the allometric intercept and exponent.

A fundamental size-based concept of EcoGENIE is the plankton cell quota for various elements. The carbon quota content (Q_c) follows a power law determining by plankton size (biovolume, *V*), scaling coefficient *a* and exponent *b* (Equation 1). And then the variable stoichiometry (Q_{1b}, Eqn. 2) is achieved by the ratio of assimilated nutrients biomass (B_{ib}, *ib* stands for P, Fe, or chlorophyll) to carbon biomass (B_c) (Droop, 1968; Flynn, 2008). This stoichiometry will limit nutrient uptake rate (Q^{stat}_{ib}, Eqn. 3) as per Geider et al., (1998), with higher value close to its maximum (Q^{max}_{ib}), lower the nutrient uptake or chlorophyll
synthesis rate . The nutrient quota range (Q^{min}_{ib}, Q^{max}_{ib}) is proportional to the carbon quota (Q_c).

$$Q_c = aV^b$$





$$\begin{split} Q_{l_b} &= \frac{B_{l_b}}{B_C}, \qquad i_b = P, Fe, Chl \\ Q_{l_b}^{\text{stat}} &= \left(\frac{Q_{l_b}^{\text{max}} - Q_{l_b}}{Q_{l_b}^{\text{max}} - Q_{l_b}^{\text{min}}}\right)^{0.1} \end{split}$$

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3.3 Plankton biomass dynamics

The biomass of any plankton group (*j*) and element (*i_b*), B_{j,i_b}, varies due to a combination of potential physiological processes that are determined by the type of organism: nutrient uptake, grazing gains, grazing losses, mortality, and respiration losses (Eqn. 5). Foraminifer-related specific processes will be introduced in following sections. We refer readers to Ward et al. (2018) for more detailed description of EcoGENIE model, although some of them might overlap.

$$\frac{\partial B_{j,ib}}{\partial t} = \underbrace{\lim_{nutrient uptake} B_{j,C} \cdot \lambda_{ib}}_{nutrient uptake} + \underbrace{B_{j,C} \cdot \lambda_{ib}}_{grazing gains} G_{j,jpecy,ib}}_{grazing gains}$$

$$\underbrace{B_{jpred,C} \cdot \sum_{jpred,I} G_{jpred,I,ib}}_{grazing losses} - \underbrace{m_j \cdot B_{j,Ib}}_{mortality loss} - \underbrace{r_{j,C} \cdot B_{j,Ib}}_{respiration loss}$$
(4)

255 3.4 Inorganic nutrient dynamics

The inorganic resource state variables (R_{i_r}) varies with nutrient uptake (V_{j,i_r}) and dissolved inorganic carbon (DIC) with the living organisms' respiration ($r_{i,c}$).

$$\frac{\partial R_{i_r}}{\partial t} = \begin{cases} \sum_{j}^{J} -\mu_{j,i_r} \cdot B_{j,c}, & i_r = Fe, P \\ \prod_{j}^{J} -\mu_{j,i_r} \cdot B_{j,c} + \sum_{j}^{J} r_{j,c}, & i_r = C \end{cases}$$
(5)

Additional sources and sinks of nutrients such as remineralisation of organic matter and air-sea gas exchange are computed in the biogeochemical module BIOGEM (Ridgwell et al., 2007).

3.4 Organic matter dynamics

Organic carbon flux (F) comes from and predators' messy feeding and the mortality loss from all plankton groups (Eqn. 6).

$$F = \sum_{\substack{j_{\text{prey}} \neq \text{trip}_{\text{pred}} = 1 \\ J}} \left(1 - \beta_{j_{\text{prey}}}\right) \left(1 - \lambda_{j_{\text{pred}}}\right) G_{j_{\text{pred}},j_{\text{pred}}} B_{j_{\text{pred}}} + \sum_{\substack{j \\ j_{\text{plankton}}}}^{J} \left(1 - \beta_{j_{\text{plankton}}}\right) m_{j_{\text{plankton}}} B_{j_{\text{plankton}}}$$
(6)

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We employ the trait-based marine plankton ecosystem model EcoGENIE (Ward et al., 2018) to simulate foraminifera's physiological processes and ecological interactions with other plankton. In this section, we summarise the core concepts specific to foraminifera modelling and refer the readers to Ward et al. (2018) for the full description of the model.⁶

In EcoGENIE, individual body size determines key physiological processes, including nutrient uptake, photosynthesis, grazing gain and predation through allometric scaling (West et al., 1997), because of its role as a master trait among pelagic organisms (Andersen et al., 2016). The modelled size-dependent parameters (except for photosynthesis) follow a generic power law: $P = aV^b$ with P the size-based parameter, V the spheric biovolume, and a, and b the allometric intercept and exponent.⁴

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 $Q_{ib} = Q_{ib}$

Metabolic processes in EcoGENIE are temperature dependent, as in the universal metabolic theory (Brown et al., 2004). Ectothermic plankton's body temperature is determined by the ambient seawater environment (T). Temperature regulation γ_T acts on metabolic processes including respiration, nutrient uptake, and predation and is modelled through an Arrhenius-like function (Eqn. 4), where the parameter A determines temperature sensitivity and reference temperature (T_{ref}) is the temperature allowing $\gamma_T = 1.\%$

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 β_{lprey} is the fraction of dissolved organic carbon (DOC) subject to advection by ocean circulation, and the remaining fraction is the particulate organic carbon (POC) subject to redistribution through the water column by sinking. Parameter β is defined by a size-based sigmoid function depending on maximum and minimum DOC fraction (β_{max} , β_{min}), and the size β_s at which DOC/POC ratio equals 1 (Ward and Follows, 2016). The proportion of DOC therefore decreases with plankton cell size.

$$\beta = \beta_{max} - \frac{\beta_{max} - \beta_{min}}{1 + \beta_s / ESD}$$
(7)

Messy feeding behaviour is modelled as unassimilated carbon fraction $(1 - \lambda_{j_{pred}})$ of prey which is limited by the sizeindependent maximum efficiency coefficient (λ_m) and the nutrient limitation (Fe or P).

$$L = \lambda_m \cdot \min_{\left[Q_P^{\text{stat}}, Q_{Fe}^{\text{stat}}\right]} \tag{8}$$

315 3.5 Plankton community structure

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The EcoGENIE plankton ecosystem model can be defined in flexible size structure. In both ForamECOGENIE 1 and 2, we resolve eight size classes of phytoplankton, seven size classes of zooplankton and one size class for each of the four foraminifer groups. Phytoplankton and zooplankton size classes include 0.6, 1.9, 6.0, 19.0, 60.0, 190.0, 600.0 and 1900.0 μm. The foraminifer cell size is set to an Equivalent Sphere Diameter (ESD) of 190 μm, which is typical for an adult foraminifer (Grigoratou et al., 2019).

<u>4 ForamEcoGENIE</u> 1, description

ForamEcoGENIE L accounted for the feeding behaviour and calcification of foraminifera (Grigoratou et al., 2019, 2021a). It implemented a predator-prey interaction (G_{jpred,jprey}, Eqn. 2) using a Holling type II model (Holling, 1965), where the overall grazing rate depends on the total available prey (F_{jpred}), the maximum grazing rate of predators (G^m_{pred}) and the half-saturation concentration of available food (k_{jprey}), and is regulated by temperature limitation (γ_T), a prey-switching term (Φ), and a prey refuge protection (1 - e^{AF_{jpred})}. The calcification trait was included by reducing foraminifera palatability (P_p which influences F_{jpred}, Eqn. 10) and mortality rate (m_j, Eqn. 5) to account for higher protection against predators and infections to the expense of a lower G^m_{foram} (Eqn. 2). To avoid repeating, we also introduce the ForamEcoGENIE 2 parameters (spine effect τ and a mixotrophy limitation λ_h) here. Readers may assume they equal one, i.e., no functioning in ForamEcoGENIE 1.

$$\begin{split} G_{j_{\text{pred}}, j_{\text{prey}}} &= \underbrace{\gamma_{\underline{T}} \cdot \underline{\lambda}_{p}}_{\text{limitations}} \cdot \underbrace{\frac{G_{pred}^{\text{m}} F_{j_{\text{pred}}}}{\mathsf{T}_{k_{j_{\text{prey}}}} + F_{j_{\text{pred}}}}}_{\text{overall grazing rate}} \cdot \underbrace{\Phi_{j_{\text{pred}}, j_{\text{prey}}}}_{\text{Switching}} \cdot \underbrace{\left(1 - e^{AF_{j_{\text{pred}}}}\right)}_{\text{prey refuge}}\right) \\ F_{j_{\text{pred}}} &= P_{p} \cdot B_{j_{\text{prey}}} \cdot \exp\left[-ln\left(\frac{\mu_{j_{\text{pred}}, j_{\text{prey}}}}{\mu_{\text{opt}}}\right)^{2}/2\sigma_{j_{\text{pred}}}^{2}\right] \end{split}$$

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(9)

(10)

In the model, the predators select their preys (Eqn. 10) based on the predator-prey size ratio $\mu_{j_{pred}, j_{prey}}$ relative to the optimal	
value μ_{opt} , predators' food range σ_{jpred}^2 , and the calcification protection P_p . For a minifera in both For a mEcoGENIE 1 and 2 were	

345 set as herbivores,
 The grazing process like other metabolisms in EcoGENIE are temperature dependent, as in the universal metabolic theory (Brown et al., 2004). Ectothermic plankton's body temperature is determined by the ambient seawater environment (*T*). Temperature regulation γ_T acts on metabolic processes including respiration, nutrient uptake, and predation and is modelled through an Arrhenius-like function (Eqn. 12), where the parameter *A* determines temperature sensitivity and reference
 350 temperature (*T*_{ref}) is the temperature allowing γ_T = 1.

 $\gamma_T = e^{A(T-T_{\rm ref})}$

The prey-switching term $(\Phi_{j_{pred}, j_{prey}})$ simulates the feeding habitat of zooplankton (Eqn. 9). The exponential *s* defines the active level of zooplankton predators, which capture abundant prey with higher priority when *s* increases. Foraminifera both in ForamEcoGENIE 1 and 2 are assumed to be ambush passive predators with s=1.

s

$$\Phi_{j_{\text{pred}},j_{\text{prey}}} = \frac{\left(F_{j_{\text{pred}}}\right)}{\sum_{j_{\text{prey}}=1}^{J} \left(F_{j_{\text{pred}}}\right)^{S}}$$
(12)

Finally, to further approach the reality, the refuge term $(1 - e^{\Lambda F_{j_{pred}}})$ in the Eqn. (2) is added to decrease grazing rate when prey availability lowers with the protection strength of Λ .

5 ForamEcoGENIE 2; improved calcification and more functional groups,

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360 <u>In ForamEcoGENIE 2</u>, we expand the first version by adding symbiosis and spine traits to the foraminifera group to distinguish four functional groups of planktic foraminifera (Table 1, Figure 1). We also implement a new calcification trade-off relating calcification energetic cost to the respiration term.

Table 1. The four modelled functional groups of planktic foraminifera and their species representative in ForamEcoGENIE 2.0.

Spine trait	Symbiosis trait	Species example	Species number*	Model implementation
Spinose	Symbiont-obligate	Globigerinoides ruber	17	This study
Spinose	Symbiont-barren	Globigerina bulloides	2	This study
Non-spinose	Symbiont-facultative	Neogloboquadrina dutertrei	5	This study

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Figure 1. Schematic diagram of ForamEcoGENIE structure, showing the biogeochemical tracers (C, Fe, PO4) in different
 colors and plankton populations with various size classes. Physiological processes here include nutrient uptake (red arrows), organic matter production caused by messy feeding and mortality (dashed arrows), and zooplankton grazing (black arrows).
 <u>A</u>. symbiont-barren spinose group; B. Symbiont-facultative non-spinose group; C. symbiont-barren non-spinose group. D. symbiont-obligate spinose group. DIC: dissolved inorganic carbon.

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5.1 Calcification trait trade-offs

5.1.1 Benefit: Mortality protection

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We modified the metabolic cost of calcification defined in Grigoratou et al. (2019, 2021a) by replacing the original reduced maximum growth rate (or specifically maximum grazing rate) with a temperature-dependent respiration loss term. We choose this new loss term over a reduced growth rate because (1) extra respiration is a more biologically realistic cost and (2) a temperature-dependent term helps reduce the low-latitude biomass as observed." *Metabolic* cost

The respiration r_j present in Eqn. 5 scales with carbon biomass and is multiplied by constant r_b and temperature limitation (Eqn. 10). We assumed that the lost carbon from respiration is instantly recycled back to DIC pool.⁴

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The mortality loss term in the generic zooplankton scales with a basal rate constant m_b (Eqn. 5). Like in Grigoratou et al. (2019, 2021a), m_b for foraminifera is downscaled by a protection term P_m where a lower value of m_j indicates a higher protection from the foraminifera test against viral and bacterial infections.

Like in ForamEcoGENIE 1.0, calcification protects from grazing and is defined by P_n , which reduces the biomass loss from

 $m_j = P_m \cdot m_b$

(13)

(14)

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420 5.1.3 Cost: higher metabolic cost

predation (Eqn. 10).

5.1.2 Benefit: Protection from predators (palatability)

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We modified the metabolic cost of calcification defined in Grigoratou et al. (2019, 2021a) by replacing the original reduced maximum growth rate (or specifically maximum grazing rate) with a temperature-dependent respiration loss term. We choose this new loss term over a reduced growth rate because (1) extra respiration is a more biologically realistic cost and (2) a temperature-dependent term helps reconcile the low-latitude biomass as observed. The respiration r_j present in Eqn. 5 scales
with carbon biomass and is multiplied by constant r_b and temperature limitation (Eqn. 11). We assumed that the lost carbon from respiration is instantly recycled back to DIC pool.

$$r_j = r_b \cdot \gamma_T$$

430 Spines are foraminifer's taxonomical basis and are made up of calcite. The biological functions of spines are related to stability in water columns, symbiosis and feeding behaviour (Schiebel and Hemleben, 2017),

5.2.1 Benefit: Enhanced grazing

Studies show that spinose foraminifera are more efficient in capturing and digesting prey thanks to the spine and rhizopodia networks (Anderson and Bé, 1976). Spines widen the prey availability of immotile foraminifer and facilitate capturing larger preys, while non-spinose species cannot hold active prey and only accept smaller particles of copepods in the laboratory observations (Anderson et al., 1979; Hemleben et al., 1989). Grigoratou et al., (2021b) modelled such benefit by reducing the half-saturation constant (conventionally noted as k in a Michaelis-Menten model). Here we adopt this approach by reducing $k_{j_{prey}}$ by a scaling parameter $\tau (0 < \tau < 1; Eqn. 10)$.

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5.2.2 Other trade-offs as calcification: higher metabolic cost and reduced palatability

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Deleted: a. Higher metabolic cost and reduced palatability⁴ We assume that the metabolic cost and protection from the spines are characterised the same way as for calcification (Eqn. 10-11), with spinose foraminifera having a higher cost and a stronger protection than non-spinoses (Table 2).⁴

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460 We assume that the metabolic cost and protection from the spines are characterised the same way as for calcification (Eqn. 13-14), with spinose foraminifera having a higher cost and a stronger protection than non-spinoses (Table 2). But there is no mortality reduction regarding spine trait.

5.3 Symbiosis trait trade-offs

465 Symbiosis is a novel trait in the model, commonly seen in marine organisms including foraminifer. Many planktic foraminifera harbour algae (e.g., dinoflagellate, diatom) within their cell (Takagi et al., 2019). We represent these symbiotic species in the model as a single organism, which combines hetero- and autotrophy, equivalent to a calcifying mixotroph. We use the trait-based representation of mixotrophy of Ward and Follows, (2016), where any plankton can "naturally" predate and photosynthesize, where the alternative strategy for specialist group is turned off (i.e., *V_m* is 0 for zooplankton and *G_m* is 0 for phytoplankton), and for mixotrophs turned on.

5.3.1 Benefit: enabled autotrophy for planktic foraminifer

The symbiont owns an independent size defined as a symbiont/foraminifera size ratio ψ (Eqn. 16) to characterise the symbiont affinity in taking up nutrients and light. Photomicrograph observations showed that foraminifera symbionts are about 1:20 smaller in size than the host cell (Takagi et al., 2019).

$$V_s = \psi^3 V_s$$

The generic nutrient uptake of symbionts follows a mixotrophy- (λ_s) , quota- $(Q_{i_T}^{\text{stat}})$ and temperature-limited (γ_T) Michaelis-Menten function where the variable (**R**) represents nutrient resources and half-saturation constant is replaced by nutrient affinity, a more mechanistic parameter for nutrient uptake α . Nutrient affinity is often referred as "clearance rate" and regarded as a proxy of competitive strength (Fiksen et al., 2013). According to Edwards et al., (2012)' s review on phytoplankton trait

 $t_{rade-offs}$, scaled nutrient affinity is negatively related to cell size because of lower surface to volume ratio, while maximum uptake rate (V_m) is positively related.

$$\mu_{i_r} = \lambda_s \cdot Q_{i_r}^{\text{stat}} \cdot \gamma_T \cdot \frac{V_{i_r}^{\text{rad}} \alpha_{i_r} \mathbf{R}}{V_{i_r}^{m} + \alpha_{i_r} \mathbf{R}}$$
(16)

The symbionts' photosynthesis growth is modelled following the size-dependent unimodal equation (Geider et al., 1998;
 Moore et al., 2001), which has shown significant explanatory power for eukaryotes phytoplankton cells than power law (Bee et al., 2008). The maximum photosynthesis rate P^m_C is determined by dimensionless parameter P_a, P_b, P_c and the biovolume of symbiont V_s, and the mixotrophy cost λ_s.

$$P_c^m = \frac{\lambda_s (P_a + \log_{10} V_s)}{P_b + P_c \log_{10} V_s + \log_{10} V_s^2}$$
(17)

The practical photosynthesis rate is further constrained by nutrient availability (the smallest between γ_{Fe} and γ_{P}), temperature 490 (γ_{T}), light intensity (γ_{I}).

$$P_{C} = P_{C}^{m} \cdot \min[\gamma_{P}, \gamma_{Fe}] \cdot \gamma_{T} \cdot \gamma_{I}$$
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The nutrient limitation γ_{i_r} (i_r is either P or Fe, see the definition in Eqn. 2) takes the minimal value from phosphorus or iron limitation term, which follows the quota relationship in Droop (1968).

Vi.

$$= \frac{1 - Q_{l_r}^{im}/Q_{l_r}}{1 - Q_{l_r}^{im}/Q_{l_r}^{mx}}, \quad i_r = Fe, P$$
(19)

As for the light limitation, it follows Moore et al., (2001) model where I represents light intensity, α is initial slope of P-I curve limited by Fe content (γ_{Fe}), and Q_{Chl} is chlorophyll quota.

$$\gamma_{l} = 1 - \exp\left(\frac{-\alpha \cdot \gamma_{Fe} \cdot Q_{Chl} \cdot l}{P_{C}^{m} \cdot \gamma_{T} \cdot \min[\gamma_{P}, \gamma_{Fe}]}\right)$$
(20)

500 5.3.2 Cost: downgrading autotroph and heterotroph efficiency,

The cost of mixotrophy is that both autotrophy and heterotrophy parts (i.e., photosynthesis and grazing rates) are scaled down (by multiplying factor λ_s and λ_h with respect to symbionts and hosts, 0 < λ_s, λ_h < 1, Eqn.9 and 16) compared to the pure auto/heterotroph specialist (Ward and Follows, 2016). Despite no direct and sufficient evidence for such a foraminifer predation cost, this assumption is common practice in mixotroph models (Castellani et al., 2013; Våge et al., 2013; Ward and 505
 Follows, 2016). We also distinguish between symbiont-obligate and symbiont-facultative foraminifera using different λ_s/λ_h parameter values to reflect their different extent of dependency on symbionts (Table 2)_x

5.4 Approximating foraminifera calcite export

Planktic foraminifera produce organic carbon in the subsurface water column (Salter et al., 2014) and sequester tests of
 inorganic carbon into the deep oceans (Schiebel, 2002). In the model the organic carbon flux was consistent as section 3.4.
 <u>The additional calcite export is approximated by multiplying the foraminifera bulk organic carbon export with a globally uniform particle inorganic carbon (PIC) to organic carbon (POC) ratio, based on the empirical first-order average value (0.36) (Schiebel and Movellan, 2012)</u>

515 6 Model experiments and evaluation

6.1 Experiments

We run an ensemble of 1,200 model experiments, each testing a different combination of parameter values, to explore all possible trait values and select the best trait combinations to match the foraminifera observations (section 3.2). The parameter sets are generated using Latin Hypercube Sampling (LHS) algorithm, uniformly sampling values of 12 model parameters

520 characteristics of foraminifer traits (Table 2; Sarrazin et al., 2016). Each simulation is run for 250 years continuing from a 10,000-year spin-up simulation as ecosystem structure typically reaches equilibrium after ~50 years. The other ecosystem parameters are the same as Ward et al. (2018) (Table S3).

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We also distinguish between symbiont-obligate and symbiontfacultative foraminifera using different λ_a/λ_h parameter values to reflect their different extend of dependency on symbionts (Table 2).[¶] We also distinguish between symbiont-obligate and symbiontfacultative foraminifera using different λ_a/λ_h parameter values to reflect their different cont of dependency on symbionts (Table 2).[¶]

a. Symbiont cell size

We determine the cell size of the symbiont from a defined symbiont/foraminifera size ratio ψ (Eqn. 12) to characterise the symbiont affinity in taking up nutrients and light. Photomicrograph observations showed that foraminifera symbionts are about 1:20 smaller in size than the host cell (Takagi et al., 2019).

b. Symbionts' inorganic nutrient uptake

The generic nutrient uptake of symbionts follows a mixotrophy- (λ_s) , quota- (Q_{ij}^{tot}) and temperature-limited (γ_T) Michaelis-Menten function where the variable (**R**) represents nutrient resources and half-saturation constant is replaced by nutrient affinity, a more mechanistic parameter for nutrient uptake α . Nutrient affinity is often referred as "clearance rate" and regarded as a proxy of competitive strength (Fiksen et al., 2013). According to Edwards et al., (2012)'s review on phytoplankton trait trade-offs, scaled nutrient affinity is negatively related to cell size because of lower surface to volume ratio, while maximum uptake rate (I_{m}) is positively related.[¶]

$V_{i} = 1$

c. Symbionts' photosynthesis¶

The symbionts' photosynthesis growth is modelled following the size-dependent unimodal equation (Geider et al., 1998; Moore ... [3]

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We use a

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Table 2 List of the foraminifer-relevant parameters tested in the global sensitivity analysis (GSA) and identified optimal

parameter values for each group,

Related trait(s)	Parameter	Description	Tested	Unit	Optimal	Optimal	Optimal	Optimal
			range [†]		values* (bn)	values* (bs)	values* (sn)	values* (ss)
	<u><i>p</i></u> _m	Protection from mortality	[0-1]		0.6	<u>0.6</u>	0.6	0.6
Calcification/spine	<u>p</u> e	Protection from palatability	[0-1]		<u>0.8</u>	<u>0.7</u>	<u>0.8</u>	0.7
	<u>r</u>	Respiration rate	[0-0.02]	mmol C d ⁻¹	<u>0.04</u>	0.06	<u>0.04</u>	0.06
Spine	<u>τ</u>	coefficient of grazing half saturation	[0-1]		<u>/</u>	<u>0.9</u>	<u>/</u>	0.9
	<u>Ψ</u>	symbiont to foraminifer size ratio	[0-0.05]		<u>/</u>	<u>/</u>	0.0015	0.0015
Symbiosis	<u>A</u> e	symbiont autotroph efficiency	[0-1]		<u>/</u>	<u>/</u>	<u>0.2</u>	<u>0.8</u>
	<u>λ</u>	foraminifer heterotroph efficiency	[0-1]		<u>/</u>	<u>L</u>	<u>0.8</u>	0.55

bn, symbiont-barren non-spinose; bs, symbiont-barren spinose; sn, symbiont-facultative non-spinose; ss, symbiont-obligate spinose

⁶⁷⁵ * For any other plankton group without these traits, scaling parameters are set to 1 and cost parameters are set to 0.
[†] GSA range is set to [0-1] for scaling parameters; respiration terms are as follow Ward et al., (2018); symbiont cell size ratio <u>upper bound</u> is calculated from observation (Takagi et al., 2019).

6.2 Comparison to observations

- 680 <u>The model generates biomass, export production for each group, and relative abundance is estimated based on export</u> <u>production</u> To validate these model outputs, we compiled 3 global adult foraminifer (>150 µm) datasets; (1) plankton net tows mostly taken from the first 100 m, (2) <u>sediment</u> core-top representing the Late Holocene (pre-industrial) and (3) seasonally resolved sediment trap time-series. We use the core-top dataset of individual abundance count to validate the spatial patterns of the relative <u>abundance</u> of each foraminifer group. The <u>plankton</u> net tow, datasets <u>are used</u> to validate the living stocks. And
- 685 the sediment trap data are used to compare with foraminifer-derived carbon export. The core-top data comes from ForCenS (Siccha and Kucera, 2017), and the plankton tow and sediment trap data from the compilation by Grigoratou, (2019) with additional foraminifera groups and sample sites. The full list of plankton tow and sediment trap data sources is in Tables S1 and S2.

For all the datasets, we firstly sum species into functional groups using species traits defined by Schiebel and Hemleben, (2017)
 and Takagi et al., (2019) (Table S4). Next the sediment trap and plankton tow data are taken annual average for spatial comparison. We also regrid the observations to be consistent to our model, which means that all the observations within a grid are averaged. Finally, we use Mielke measure (next section) to quantify the model-data fit. Note (1) this is for spatial comparison and therefore we do not take annual average and calculate skill scores for seasonal model-data comparison (2) the

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Arctic and the Mediterranean Sea are omitted in comparison because of the resolution-derived model's poor ocean physics in those regions,

715 The units of plankton net tow ("count m³") and sediment trap data ("count m² d⁻¹") are converted into "mmol C m³" and "mmol C m² d⁻¹" using the empirical factor of 0.845 µg count¹ from Schiebel and Movellan (2012). The model generates carbon export production in "mmol C m³ d⁻¹", so it is transformed to "mmol C m² d⁻¹" by multiplying the surface layer depth (80.8 m). Note that although the sediment traps were deployed in different depths (averagely 1.96 km) and seasons. we ignore the vertical degradation because foraminifer tests are relatively fast-sinking due to large size (Takahashi and Be, 1984) and we

720 do not have well-built attenuation profile for foraminifer tests,

We considered species with less than 3% abundance as absent to avoid the uncertainty caused by ocean currents transportation (van Sebille et al., 2015). We applied this threshold based on the standard error of fisher's α diversity index (Fisher et al., 1943),

725 6.3 Performance metrics

We use, Mielke measure, or M-score (Watterson, 1996; Watterson et al., 2014) to quantify the model-data fitting goodness (Eqn. 21). This metric is essentially non-dimensional transformed mean square error, and has been widely used in general circulation models (Gregoire et al., 2011, Anon, 2016). This score spans from -1 to 1 with values closer to 1 representing better model performance, 0 stands for no predicting skills. And negative M-score indicates negative correlation.

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$$M = \frac{2}{\pi} \arcsin\left[\frac{\sum_{i=1}^{n} (M_i - O_i)^2 / n}{\sigma_m^2 + \sigma_o^2 + (\mu_m - \mu_o)^2}\right]$$

The numerator is mean square error, with M_i and O_i the model and observational value in the *i*th grid, and *n* the total number of grids. σ^2 and μ are the variance and mean, with superscripts *m* and *o* representing model and observed fields, respectively.

6.4 Global Sensitivity analysis

35	We conduct a global sensitivity analysis (GSA) to explore the model robustness of our 1,200 experiments using the PAWN
	method (Pianosi and Wagener, 2015). This method measures the sensitivity of model outputs (focusing on M-score here) to
	different values of input parameters. A total M-score is calculated by summing scores of each foraminifer group in biomass,
	POC export, and relative abundance (i.e., the total score ranges from -12 to 12). To further measure the uncertainty and
	robustness of the GSA results, we also apply a bootstrapping method with 1,000 resamples. This method allows us to
40	understand the confidence intervals of the sensitivity indices without running more experiments (Wagener and Pianosi, 2019)
	We bootstrapped our data using <i>rsample</i> package (Silge et al., 2021) in the R software environment v4.1 (R Core Team, 2021).

7 Results and discussions

7.1 Model ensemble results

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Deleted: To further constrain the uncertainty of the observational data, we applied median absolute deviation (MAD) measurement to detect the most robust and close-to-reality data. Finally, we grouped the species-based data into functional groups (Table 1) using species traits defined by Schiebel and Hemleben, (2017) and Takagi et al., (2019) (Table S2).

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We also perform basic statistics (mean, standard error, sum) on the model and observational data. We use the standard error of the mean to represent the accuracy of the sampling mean, particularly for the observational studies. Finally, we do not consider model outputs for the Arctic and the Mediterranean Sea because of limitations with occan physics due to low model resolution in that region.

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Overall, the 1,200 model runs fit generally reproduce POC export (<u>summed M-score: $-0.2_{\sqrt{2}} \sim 1.0$) and relative abundance</u> (<u>summed M-score: $-0.5_{\sqrt{2}} \sim 1.2$ </u>), with poorer comparison with plankton net data (<u>summed M-score: $-0.1_{\sqrt{2}} \sim 0.5$ </u>) (Figure 2). A

- 780 heatmap of M-Scores (Figure 2) shows the experiments cluster into 4 groups with respect to the 3 observation metrics. Most parameter combinations (cluster C) show low performance in predicting foraminifer metrics. Cluster D shows an inverse relative abundance distribution to the observation, while cluster B has only good POC export performance. Cluster A which includes our optimal parameters, achieves the highest (best) abundance M-score and has good biomass and POC export prediction. We notice that this cluster accompanies low foraminifer export production, which indicates that the higher-than-
- 785 observation biomass/carbon are the primary constrain for the model to get better scores. The sensitivity analysis result confirms this. Our model performance is mostly sensitive to the symbiosis efficiency (λ_5), and the palatability protection (P_p) from both spine and calcification which both directly govern the nutrient uptake and loss (Figure 3). Models with low export production and higher scores associate with <200 µm foraminifer size and small symbiont size (symbiont/host size ratio <0.02) which facilitate dwelling in the tropical regions, realistic shell protections and high respiration cost that do not lead to striking
- 790 biomass/export (Figure S3).

We select the <u>optimal parameter set</u> with the highest total M-score score (<u>Table 3</u>), which also has the highest M-score for the relative abundance (group mean = $0_{\frac{1}{2}}$) and POC <u>export</u> (group mean = $0_{\frac{1}{2}}$). In this run, all the foraminifera groups achieve the highest total M-score (Figure S2). More details about this optimal model is shown in next sections.

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incorporating all four main foraminiferal ecogroups in the modern ocean without weakening the overall ability of predicting ecosystem body size, biomass and POC export.

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Figure 2. M-score heatmap of the model ensemble compared with foraminifera "Biomass" (plankton net data), "POC export" (sediment trap data), groups' <u>"Relative</u> Abundance" (sediment core top data). Each column sums the M-score of four foraminifer groups, and the fourth column is the sum of all the left three. The right panel shows the global annual mean export production of all foraminifer groups. The ensemble cluster was derived from a complete linkage clustering algorithm (Jarman, 2020). The higher the M-score value the better the performance, negative value stands for negative correlation.

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sediment trap and plankton net studies (Figure 4.5). The symbiont-obligate spinose group is the most abundant group with a global relative abundance of 60.7% (M-score: 0.32, Figure 4g) dominating in the tropical open oceans, while the symbiont-barren non-spinose (Figure 4a) and spinose groups (Figure 4c) both mainly occupy the high latitudes (M-score = 0.51 and

940 0.31, respectively). The symbiont-facultative group (Figure 4e) is underestimated with the visible model-data disparities in the eastern equatorial pacific where the sediment core shows exclusively high abundance. This may be due to the resistant dissolution of some species' test (e.g., *N. dutertrei*) as suggested in a previous model study (Lombard et al., 2011). Overall, the Root Mean Square Error (RMSE) of relative abundance in this model (12% to 42%, Table \$5) is comparable to the species-based models, like FORAMCLIM (5-23%, Lombard et al., 2011) and PLAFOM (22-25%, Fraile et al., 2009). This indicates

945 that symbionts and spines are sufficient to explain the variance of geographic distribution. <u>While non-symbiotic foraminifer follows the biogeography of prey in high nutrient areas, the symbiont size (0.0015 symbiont/foraminifer) indicates that the small body size and therefore high nutrient affinity is a key factor for symbiotic foraminifer in low-nutrient areas like subtropical gyres. Although this actual symbiont size does not necessarily be around 0.0015 times foraminifer size, the fact that foraminifer can host up to 3,000 algal cells can explain the high nutrient uptake</u>

950 efficiency (Spero and Parker, 1985). The symbiont-barren spinose foraminifer (Figure 4d) displays higher abundance in Arabian Sea, South China Sea than non-spinose foraminifer (Figure 4b). This might be related to the carnivory feeding (rather than enhanced grazing efficiency) of spinose foraminifer that is not included in current model.

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Figure 4. Relative abundance of the modelled geographical distribution (left column) of the four planktic foraminifer function 965 groups, compared to the ForCenS core-top dataset (right column; Siccha and Kucera, 2017). Subplots titles are the M-scores derived relative to observation and the global mean of relative abundance,

7.3 <u>Annual average b</u>iomass of foraminifera groups

As the plankton tow data, the model reproduces the low biomass of planktic foraminifer (Figure 5). The global annual mean

of biomass ranges from 0.001 to 0.010 mmol C m⁻³, or 0.08-0.8 mmol C m⁻², with the most contribution from symbiont barren 970 groups (Figure 7). Globally integrating all foraminifer groups, the model estimates 0.0068 Gt C or 6.83 Tg C biomass (Figure 7). For comparison, the MAREDAT project estimated the global average biomass of 0.24-0.94 mmol C m² which integrates



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flux rates.



all the depth from 0 to 2,500 m (Schiebel and Movellan, 2012). Schiebel and Movellan (2012) also estimated a total of 0.94~3.63 Tg C living foraminifer at any time point. Given the similar mean biomass range estimation, the difference of total biomass between our model and this data is likely caused by the grid area estimation. Therefore, it is safe to state that the model produces good estimation compared with the observation.
 The skill score, however, does not capture this good mode-data fit. This is mostly caused by regridding the data points into model grid resolution. The plankton net data are spatially concentrated in North Atlantic, Northwestern Pacific, Arabian Sea, and Indian sector of Southern Ocean. Under such circumstance, regriding causes sparser data and makes skill score sensitive

995 to several outlier grids. Therefore, the insufficient data is the major reason of low scoring in biomass.

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015 main functional groups of planktonic foraminifera.

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7.4 Annual average POC and calcite export of foraminifera groups

The model reproduces consistent geographical distributions and magnitude of POC export compared to sediment trap samples for all 4 foraminifer groups (Figure 6, 7). The model estimates $0.002-0.031 \text{ mmol C} \text{ m}^{-2} \text{ d}^{-1}$, compared with 0.001-0.026 shown in sediment trap studies. However, the total M-score for the model POC export is 0.63, not reflecting the good matchiness

between model and observation. This is similarly caused by limited data as previous biomass section.

(a) symbiont-barren non-spinose 2.90E-02
(c) symbiont-facultative non-spinose 2.44E-02
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best (0.28) and the symbiont-facultative group performing the [14]

In terms of global estimation, the model suggests the <u>foraminifer-derived</u> organic carbon export of 0.1 Gt C yr⁻¹. Symbiontbarren non-spinose taxa dominate this carbon export (55%) followed by symbiont-barren spinose, symbiont-facultative and symbiont-obligate groups (30%, 3%, and 11%, respectively). Integrating all the ecogroups and using an empirically averaged <u>PIC:POC ratio of 0.36</u> (Schiebel and Movellan, 2012), our model estimates a total calcite flux of pelagic foraminifera of 0.033 Gt <u>PIC yr⁻¹</u> (Figure 8), which is at least five times smaller to the 0.16–0.39 Gt <u>PIC</u> yr⁻¹ of Schiebel (2002) within the top 100m ocean. The reason is likely we calibrate the model with sediment traps deployed in averagely 2 km water, while the high

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flux signal in the top 300 m is not captured. In addition, the high temporal variability in the observation is also not well captured by the model.



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Detected. The character export me index ratio within the tow targe of previous estimates (1.3-3.2 Gf yr²; Schöel et al., 2001) and contributes 19% to the global marine CaCO₃ production (Milliman et al., 1999). This estimate is similar to 21% reported in Kiss et al., (2021) based on sediment traps at Cape Blanc and to Salmon et al., (2015) data from the Sargasso Sea ranging between 0.40% (but mostly < 25%). Regionally higher contributions (32-49%) have been reported in the Southern Ocean (Salter et al., 2014) who included deep-dwelling species which are not represented in this model. To summarise, our estimation of foraminifer calcite export is generally trustable to previous observational studies.

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Into set of parameters does not increase the overall ecosystem biomass or POC export, and slightly increases the mean cell size by ~0.5 µm (Figure 3c). Therefore, the incorporation of symbiosis and spines into our trait-based model successfully widens the ability of the model to represent foraminifers in the surface ocean by incorporating all four main foraminiferal ecogroups in the m (... [15])



shown in Figure 11. In addition, the cool/temperate group's earlier-when-warmer signature is also well reproduced in our model from high to low latitude. However, some detailed patterns like the occasionally two cycle per year is not shown in the model. Overall, the model reproduce the seasonal peak time well for different groups, despite the weaker amplitude than the

165 sediment traps suggest.

The foraminifer intra-annual biomass/export variability are driven by the seasonal environmental changes which determine how optimal foraminifera are in the ecological niche. Usually, temperature is assumed as the primary driver. However, the correlated factors like primary productivity are likely to be the true driver (Jonkers and Kučera, 2015). This might be the reason

170 why our model cannot produce as high production as the sediment traps because the upwelling regions in GENIE model are relatively poor constrained. In addition to the model, the low spatial and temporal resolution of observed data also limits us to have a broader overview of seasonal variation, especially in the term of biomass. We also suggest that more functional trait studies regarding temporal variability such as dwelling depth, life history cycle should help resolve this gap in the future.



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An additional source of uncertainty is the data quality, as ideally models should be calibrated against spatially and temporally abundant and well-constrained data. Therefore, the comparison of our model against plankton tow data was limited by the low temporal resolution of the data (Figure 9). The data is not only undersampled in limited regions (North Atlantic Ocean, Caribbean, Sea Arabian Sea) but also biased towards specific sampling seasons, generated using different mesh size and water depth as discussed in the first application of ForamEcoGENIE (Grigoratou et al., 2021a)

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Parameter sensitivity test

We conducted a sensitivity analysis to determine which parameters influence the model performance most. The results show that the model performance is mostly sensitive to the symbiosis trait (λ_0), and the palatability protection (P_p) from both spine and calcification (Figure 10), confirming the important role of symbiosis and protection from predators in foraminiferal ecology and the p(...[16])





sites (shown in the map with corresponding letter). Sites are selected to according to the number of comparable data points and ocean basins.

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245 Although ForamEcoGENIE 2 development focuses on the marine ecology, it also lays the foundation to future linking to biogeochemical cycles. For example, spinose foraminifer has more calcite export than non-spinose foraminifer and spines

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We conducted a sensitivity analysis to determine which parameters influence the model performance most. The results show that the model performance is mostly sensitive to the symbiosis trait (λ_i), and the palatability protection (P_p) from both spine and calcification (Figure 10), confirming the important role of symbiosis and protection from predators in foraminiferal ecology and the possibility that overestimated biomass of non-spinose foraminifer is influenced by the palatability parameterisation.⁴



Figure 11. Sensitivity of parameters to overall model performance (as combined M-scores). Bar boundaries indicate the 95-percent confidence interval with the thick line showing the mean value. Grey area indicates non-influential range of index value as control group, s, symbiont-facultative non-spinose; ss, symbiont-bearing spinose.

265 control the sinking velocity of foraminifer shell (Takahashi and Be, 1984). Symbiotic groups expand the global niche for foraminifer. The community compositional change of these groups under different climatology would influence the ocean carbon cycles by changing the calcite export and surface alkalinity. So far, no Earth system models have explicit foraminifer group participating the inorganic carbon cycle. Therefore, this would be an important direction of our model development in the future.



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9 Model limitation

Our model implementation has limitation due to our current gap in understanding trait. One is the not explicit definition of spine and symbiosis. For example, with the definition in Mitra et al., (2016), the current modelling approach falls within

Figure 12. Tuned ForamEcoGENIE 2 (third column, with four foraminifer groups) comparison with ECOGENIE (first column) and ForamEcoGENIE 1 (second column, with non-spinose non-symbiont foraminifer only) in terms of ecosystem mean size, ecosystem biomass, ecosystem POC export, and Net Primary Production (NPP). The first column displays absolute values, whilst the latter two are the ratio relative to the first column.

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constitutive (inherent or innate) mixotroph rather than foraminifer's specialist non-constitutive mixotroph. Such indirect
 photosymbiotic relationships in the model could miss the individual climate sensitivities of symbiont and host. Meanwhile, the implicit definition of spines and symbiosis cannot lead to a changeable spine and symbiosis effects with environmental changes. And this is because we lack such observational studies.

Some detailed trait interaction and variation are not included as well. The symbiotic spinose foraminifer can use their spines
 to place algal symbionts in the daytime (LeKieffre et al., 2018), which theoretically increases photosynthesis efficiency due to an increase in surface area relative to non-spinose species. And the unclear relationship between body size and spine and symbiosis forces the model the fix the foraminifer size at 190 µm and ignore the biomass of juvenile foraminifer. From this perspective, our foraminifer model belongs to a functional-type model rather than a strict trait-based model. However, we also note this study focuses on solving the gap of modelling past time and resemble the functional biogeography of planktic foraminifer.

In addition to model implementation, the spatial or temporal sparseness of plankton net and sediment trap data make modeldata comparison difficult. The M-score cannot reflect the model-data difference properly as there are limited grid points after annually averaging and regridding. The parameterisation is therefore highly influenced. The sediment trap data are temporally more continuous, but we still lack a vertical profile in different locations as most POC flux studies (Henson et al., 2012). The current model parameterisation for export production is compared with vertically averaged values. Therefore, the foraminifera surface carbon export should be higher than the model predicts and agree with the high proportion in Scheibel (2002). The reason behind observed low biomass and high carbon export is unclear but likely caused by the underestimated sampling. Because plankton net is temporally low-resolution and cannot capture the peak production signal.

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10 Ecosystem model implementation and complexity

The current coupled Earth system models are mostly NPZD (Nutrient-Plankton-Zooplankton-Detritus) based (Keller et al., 2012; Watanabe et al., 2011), or PFT (Plankton Functional Type) based (Moore et al., 2001; Aumont et al., 2015). While
 NPZD models focus on the biogeochemical fluxes and have ignored different size classes and types within phytoplankton and zooplankton models, PFT models have more explicit functional types (e.g., diatom, coccolithophore) and size classes (e.g., picoplankton, nanoplankton, microplankton). This usually helps achieve better model performance in region scale. For example, the inclusion of diatom-diazotroph assemblages allows model to predict peak production in oligotrophic areas

(Tréguer et al., 2018). Trait-based models as another member like Darwin (Follows et al., 2007) and EcoGENIE (Ward et al., 2018), solve the plankton diversity based on allometric rule. Therefore, the plankton size spectrum can be more continuous and physiological parameters can be calculated based on individual size. This is theoretically more robust especially when the model is used to simulate past ecology or large-scale pattern. However, these models also classify species into various

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The limitations in our current trait understanding still influence our model definition and performance. For example, symbiotic spinose foraminifer can use their spines to place algal symbionts in the daytime (LeKieffre et al., 2018), which theoretically increases photosynthesis efficiency due to an increase in surface area relative to non-spinose species. In this model, we do not explicitly model the photosymbiotic relationships which could be sensitive to individual climate sensitivities of symbiont and host. Similarly, spinose foraminifer arefer large zooplankton prey over phytoplankton as their prey, while non-spinose species are often herbivorous (Anderson et al., 1979). Such specialised behaviour is not resolved in the model and might explain why the symbiont-barren spinose G *bulloides* do not show opportunistic behaviour, i.e., appear earlier than other groups in the seasonal succession (Taylor et al., 2001).

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functional types according to evolutionary similarities. This might be because we cannot distinguish the ocean diversity by size trait only. More functional traits beyond size like symbiosis (Suggett et al., 2017) and body extension (Ohman, 2019) also play an important role in influencing plankton feeding, metabolism and export efficiency.

335 There is also a debate on whether higher ecosystem complexity is needed (See introduction in Quéré et al., 2005) because more parameters introduce more freedom, longer run time and the ocean physics influence biogeochemistry more than ecology. However, recent studies suggested the important role of biodiversity in the ocean biological pumps (Tréguer et al., 2018). The presence of functional group like diazotroph can even diverge the response trend to global warming (Bopp et al., 2022). Therefore, relative to the simple food web structure in current generation of models, ecosystem implementation is very likely the source of uncertainties for modelled ocean carbon storage in the future (Wilson et al., 2022). In addition, the recent model advances already included more supports to ecosystem complexities such as more functional types, variable stoichiometry, and nutrient co-limitations (Séférian et al., 2020). To our study_#the significance of more functional traits also enables us to track the foraminifer functional diversity change and provide the likelihood of explicitly linking foraminifer ecology to carbonate pump in the future.

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11 Summary

In this study, we extend the trait-based planktic foraminifer model, ForamEcoGENIE, to include symbiosis and spine traits and resolve all main foraminifer functional groups. Using Latin Hypercube Sampling, we generated 1,200 parameter samples and compared these with three global observational sources: core-top, plankton tow, sediment trap. We assessed the model

- 350 performance of biogeographic distributions, the carbon biomass and <u>foraminifer-derived</u> carbon export. Our global sensitivity analysis shows the symbiosis trait and the palatability protection of spine and test strongly influence model performance. <u>Our</u> best set of model parameters successfully reproduces the modern biogeographical distribution of the main four foraminifera ecogroups and produces a global <u>annual</u> mean <u>biomass (0.001 to 0.010 mmol C m⁻³) and foraminifer-derived</u> organic carbon export (0.002-0.031 mmol C m⁻² d⁻¹) similar to observations. The two symbiont-barren groups account for \$5% of standing
- 355 stocks and <u>foraminifer-derived</u> carbon export, while the two symbiotic group contribute the remaining <u>15%</u>. The model accurately reproduces <u>peak time of</u> seasonal time-series observations of foraminiferal biomass and organic carbon flux <u>but</u> performs poorly <u>in seasonal amplitudes</u>, <u>particularly in upwelling regions</u>. These results provide confidence in the model's ability to explore foraminifer ecology and diversity in the geological record and to interpret and question the foraminifer microfossil records, for example of the last glacial maximum, as well as helping to solve riddles about their ecology in the
- 1360 past. The trait-based framework of cGENIE ecosystem also provides potential to extend the model by presenting more traits such as life history and differential calcification rates across groups.

Code and data availability

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Deleted: The model suggests a foraminifera calcite export rate of 1.1 Gt yr⁻¹, equivalent to 19% of the global marine calcium carbonate budget. The value agrees with the lower end of modern estimates.



The source codes <u>and data</u> are archived in <u>https://doi.org/10.5281/zenodo.6808760</u>, Experiment configuration and observational dataset can be found in genie-userconfig/MS/yingetal.GMD.2022. <u>More general manual for cGENIE model can be found in https://github.com/derpycode/muffindoc.</u>

Author contribution

RY, FMM, JDW, DNS designed the study. RY, FMM, JDW developed the model code. RY performed the experiments, data collection and visualisation. All authors wrote and edited the original draft.

Competing interests

The authors declare that they have no conflict of interest.

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