Isoprene and monoterpene simulations using the chemistry-climate model EMAC (v2.55) with interactive vegetation from LPJ-GUESS (v4.0)

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Abstract.

Earth system models (ESMs) integrate previously separate models of the ocean, atmosphere and vegetation in one comprehensive modelling system enabling the investigation of interactions between different components of the Earth system. Global isoprene and monoterpene emissions from terrestrial vegetation, which represents the most important source of VOCs in the Earth system, need to be included in global and regional chemical transport models given their major chemical impacts on the atmosphere. Due to the feedbacks of vegetation activity involving interactions with the weather and climate, a coupled modelling system between vegetation and atmospheric chemistry is a recommended tool to address the fate of biogenic volatile organic compounds (BVOCs). In this work, we present further development in linking LPJ-GUESS, a global dynamic vegetation model, to the atmospheric chemistry-enabled atmosphere-ocean general circulation model EMAC. We evaluate terrestrial BVOC emission estimates from the submodels ONEMIS and MEGAN in EMAC with (1) prescribed climatological vegetation boundary conditions at the land-atmosphere interface; and (2) dynamic vegetation states calculated in LPJ-GUESS (replacing the "offline" vegetation inputs). LPJ-GUESS-driven global emission estimates for isoprene and monoterpenes from the submodel ONEMIS were found to be 546 Tg yr^{-1} and 102 Tg yr^{-1}, respectively. MEGAN prescribed 657 Tg and 55 Tg of isoprene and monoterpene emissions annually. We also evaluated the sensitivity of the new coupled system in doubling CO_2 scenarios. This work provides evidence that the new coupled model yields suitable estimates for global BVOC emissions that are responsive to vegetation dynamics. We conclude that the proposed model setup is a useful tool for studying land-biosphere-atmosphere interactions in the Earth system.

1 Introduction

The land surface of the Earth is dominated by vegetation, with forests covering ∼42 million km^2 in tropical, temperate and boreal regions, making up ∼30% of the total land area (Bonan, 2008). The terrestrial biosphere is known to be a primary source of volatile organic compounds (VOCs) such as isoprene and various terpenes, accounting for around 90% of the total VOC emissions to the atmosphere (Guenther et al., 1995). The processes driving VOC emissions from plants are complex and not

fully understood - VOCs seem to play a role in protecting photosynthetic activity in plants from damage caused by reactive oxygen species, which are synthesised in leaves at high temperatures (Niinemets, 2010; Harrison et al., 2013; Lantz et al.,

25 2019). Not all plant species emit VOCs and this adaptation can be considered as an evolutionary trait that benefits certain plant species in hot, dry environments (Taylor et al., 2018). Isoprene and monoterpenes are the most abundant species among the biogenic volatile organic compounds (BVOCs) (Kesselmeier and Staudt, 1999; Lathiere et al., 2006; Guenther et al., 2012), and their high reactivity exerts a significant influence on atmospheric composition (Atkinson, 2000). The atmospheric chemical lifetime of such BVOCs ranges from minutes to hours (Atkinson and Arey, 2003) as they quickly interact with tropospheric

30 species including carbon monoxide (CO), hydroxyl radicals (OH) and ozone (Lelieveld et al., 1998; Granier et al., 2000; Poisson et al., 2000; Pfister et al., 2008), thus altering the atmosphere's oxidation capacity. BVOCs are also the main precursor for secondary organic aerosols (SOA), which can exert a significant forcing on the radiative balance of Earth, both directly through scattering and absorption of solar radiation, and indirectly through changing cloud properties (Rap et al., 2013; Scott et al., 2014). SOA also contributes to change in the radiation balance by decreasing the solar near-surface direct radiation while

35 at the same time increasing the diffusive radiation contribution (Wang et al., 2019).

The first BVOC models employed empirical relations describing isoprene emission rate dependencies on temperature and light (Guenther et al., 1991, 1993; Tingey et al., 1981), and montoerpene emission rate dependency on temperature (Evans et al., 1982; Lamb et al., 1987; Tingey et al., 1980, 1981). The formulations include a species or vegetation type-specific emission factor characterised from field or laboratory measurements (e.g. Lamb et al., 1985; Arey et al., 1991; Guenther et al.,

40 1993) which is defined for arbitrarily chosen environmental conditions (usually $30°C$ and 1000 μmol photons m^{-2} s^{-1}) (Grote and Niinemets, 2007). This approach has been extensively used to study different ecosystem types all over the world including desert (Geron et al., 2006), grassland (Bai et al., 2006), savanna (Guenther et al., 1999; Otter et al., 2003), Mediterranean (Cortinovis et al., 2005), tropical (Harley et al., 2004), temperate (Karl et al., 2003) and boreal forests (Westberg et al., 2000). Empirical algorithms are also presently used in well-established global BVOC models such as ONEMIS (Kerkweg et al.,

45 2006) and MEGAN (Guenther et al., 2006). Another category of models, know as process-based models, have been derived from knowledge of biochemical processes rather then purely empirical relations (e.g. Niinemets et al., 1999; Bäck et al., 2005; Arneth et al., 2007b; Schurgers et al., 2009). Such models also consider emission inhibition from carbon and energy availability, allowing for stress effects to be taken into consideration. A recent study proposes an new approach where VOC emissions can be modelled via the strong linear dependency between isoprene emissions and the leaf's 'energetic status' (Harrison et al.,

50 2021).

The land-biosphere-atmosphere interface in models is fundamentally important to study the fate of BVOCs in the atmosphere, yet, early models were designed to simulate single components of the Earth system in isolation, prescribing simple non-interacting boundary conditions at the interface. However, models have become increasingly coupled with dynamic multidirectional fluxes between the different models considered. This yielded a new category of models we now call Earth System

55 Models (ESMs). ESMs are highly comprehensive tools ideal for modelling past and future climate change with biogeochemical feedbacks and also for studying biosphere-atmosphere interactions explicitly (Flato et al., 2014). To this end, several modelling studies have linked atmospheric chemistry-enabled models with dynamic vegetation models to investigate the im-

pacts of changing vegetation cover on global atmospheric emissions, atmospheric chemistry, and future climate change (e.g. Levis et al., 2003; Sanderson et al., 2003; Naik et al., 2004; Lathiere et al., 2005; Arneth et al., 2007b). Empirical-based models

60 suggest increased BVOC emissions in future climate scenarios resulting from temperature sensitivity and enhanced vegetation activity in a CO_2-richer atmosphere (Sanderson et al., 2003; Naik et al., 2004; Lathiere et al., 2005). In contrast, process-based models indicate that CO_2-inhibition of leaf-isoprene metabolism can be large enough to offset increases in emissions (Arneth et al., 2007b; Heald et al., 2009). More recent laboratory studies provide evidence that certain plant species emit more isoprene in high CO_2 environments (e.g. Sun et al., 2013), challenging the significance of CO_2 inhibition effects. A modelling study

65 also showed that regardless of whether CO_2 inhibitory effects are considered or not, the temperature is the most important driver for increased isoprene emissions (Cao et al., 2021).

Sporre et al. (2019) employed an ESM to investigate climate forcing caused by BVOC-aerosol feedbacks, where it was determined that increased BVOC emissions and subsequent SOA formation in future climate scenarios result in -0.43 W m^{-2} stronger net cloud forcing and -0.06 W m^{-2} forcing from direct scattering of sunlight. A new ESM that integrates

70 the chemistry-climate model EMAC (Roeckner et al., 2006; Jöckel et al., 2005) with the dynamic global vegetation model (DGVM) LPJ-GUESS (Smith et al., 2001; Sitch et al., 2003; Smith et al., 2014) has been recently developed (Forrest et al., 2020). In a first study, the coupled model gave a good representation of worldwide potential natural vegetation distribution, despite some regional variations, especially at lower spatial resolutions. In this study, we present further model coupling of LPJ-GUESS within the EMAC modelling system with respect to vegetation driven emissions. We evaluate global isoprene and

75 monoterpene emission fluxes from the submodels ONEMIS and MEGAN in EMAC with online dynamic vegetation inputs derived from LPJ-GUESS. We additionally analyse temperature and fertilisation effects in doubling CO_2 climate scenarios.

2 Methods

2.1 EMAC modelling system (v2.55)

The EMAC (ECHAM/MESSy Atmospheric Chemistry) model is a numerical chemistry and climate modelling system that

80 includes submodels that describe tropospheric and middle atmosphere processes, as well as their interactions with oceans, land, and anthropogenic activities. It originally combined the ECHAM atmospheric GCM (Roeckner et al., 2006) with the Modular Earth Submodel System (MESSy) (Jöckel et al., 2005) framework and philosophy where physical processes and most of the infrastructure has been divided into "*modules*", which can be further developed to improve existing process representations and new modules can be added to represent new or alternative process representations. In recent years EMAC has been further

85 developed to include a broader representation of atmospheric chemistry by coupling different processes such as representations for aerosols, aerosol–radiation and aerosol–cloud interactions, e.g. Tost (2017). In this study, the version 2.55 has been utilised, which is based on the well documented version used in comprehensive model intercomparison studies, e.g. Jöckel et al. (2016).

BVOC modules in EMAC

Both ONEMIS and MEGAN are emission modules which are based on the Guenther algorithms (Guenther et al., 1993, 1995).

90 In this schemes emissions are calculated as a function to ecosystem specific emission factors, surface radiation, temperature,

the foliar density and its vertical distribution. The schemes mostly differ in the evaluation of canopy process for light and temperature sensitivity. Technical details for canopy processes employed in ONEMIS can be found in Ganzeveld et al. (2002), while Guenther et al. (2006) provides details for MEGAN. Here we use the parameterized canopy environment emission activity algorithm in MEGAN rather then the alternative detailed canopy environment model that calculates light and temperature

95   at each canopy depth.

2.2 LPJ-GUESS DGVM (v4.0)

Lund-Potsdam-Jena General Ecosystem Simulator (LPJ-GUESS) (Smith et al., 2001; Sitch et al., 2003; Smith et al., 2014) is a DGVM featuring an individual-based model of vegetation dynamics. These dynamics are simulated as the emergent outcome of growth and competition for light, space and soil resources among woody plant individuals and a herbaceous understorey in

100   each of a number (50 in this study) of replicate patches representing random samples of each simulated locality or grid cell. The simulated plants are classified into one of a twelve plant functional types (PFTs) discriminated by growth form, phenology, photosynthetic pathway (C3 or C4), bioclimatic limits for establishment and survival and, for woody PFTs, allometry and life history strategy. LPJ-GUESS has already been implemented in global ESMs (e.g. Weiss et al., 2014; Alessandri et al., 2017), and more recently coupled with EMAC (Forrest et al., 2020). Additionally, LPJ-GUESS has been widely assessed and

105   also extended to include several terrestrial processes leading to over 450 International Scientific Indexing publications[1]. LPJ-GUESS currently gives potential natural vegetation rather than present day vegetation, hence the current configuration cannot be validated yet. However, land use schemes will be included in a future version of LPJ-GUESS, allowing for a more realistic representation of the vegetation dynamics.

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[1]https://web.nateko.lu.se/lpj-guess/LPJ-GUESS_bibliography.pdf (last access: 14 April 2022)

2.3 LPJ-GUESS-EMAC coupling for BVOC emission estimates

2.3.1 Overview of the coupled setup
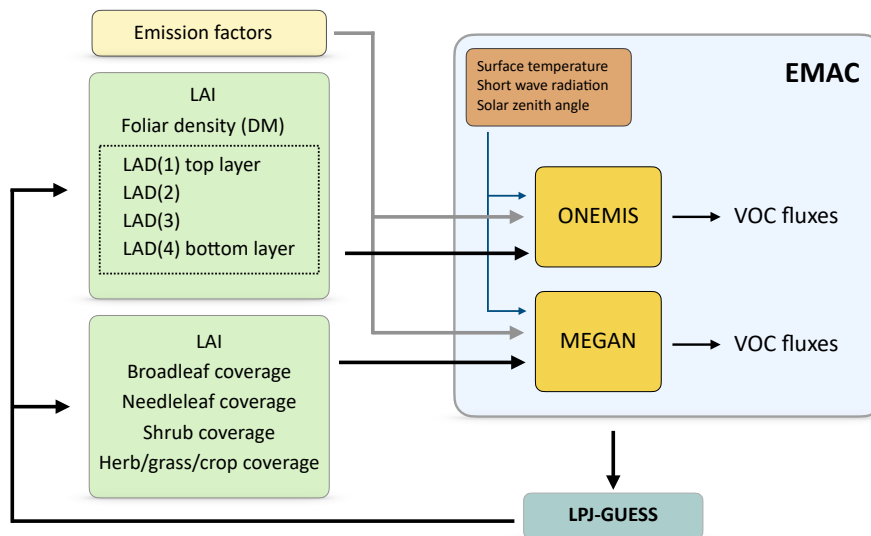
Figure 1. Model setup for BVOC emission estimates in EMAC. The vegetation variables needed in ONEMIS and MEGAN are now provided by LPJ-GUESS, replacing offline climatological datasets.

In this work, we extend the model coupling strategy employed in Forrest et al. (2020), where modifications are done in LPJ-GUESS such that it provides its functionality via a new submodel in the MESSy framework, yet keeping the LPJ-GUESS source code intact with minimal modifications. At the end of the simulation day, EMAC provides LPJ-GUESS the daily-mean 2m temperature, daily-mean net downwards shortwave radiation, and the total daily precipitation. Daily CO_2 concentrations and nitrogen deposition are also given by EMAC to LPJ-GUESS. As a result, the LPJ-GUESS land surface conditions are entirely determined by the EMAC atmospheric state and chemical fluxes. In Forrest et al. (2020), only one-way coupling was performed, which means that the land surface vegetation condition calculated in LPJ-GUESS has no effect on the atmospheric state in EMAC. In this study, we leverage some of the existing LPJ-GUESS output variables to calculate isoprene and monoterpene emission rates in EMAC, allowing for the first time a limited two-way coupling between the two modelling systems.

Fig.1 illustrates the model configuration for computing isoprene and monoterpene emissions fluxes in EMAC using the submodels ONEMIS and MEGAN. Both models require emission factors for the various PFTs, the solar zenith angle, surface radiation, and surface temperature. Additionally, ONEMIS requires the *vegetation* variables; leaf area index (LAI), foliar density, and the leaf area density (LAD) canopy profile, while MEGAN needs the LAI and fractional coverage of broadleaf, needleleaf, grass and shrub ecosystem types. In the original setup, the vegetation input variables are prescribed from offline climatological datasets, whereas, in the new configuration, we replace the climatological vegetation variables with ones calculated

online in LPJ-GUESS. This implies that the new setup feeds dynamic vegetation states to the BVOC modules that are directly computed in LPJ-GUESS on a daily time-scale and driven by atmospheric states and chemical fluxes in EMAC, allowing for estimates of isoprene and monoterpene emissions with dynamic vegetation. In this work, we adopt new calculations from the LPJ-GUESS output to derive all input vegetation variables needed for the calculations of BVOC emission fluxes in EMAC via

130 ONEMIS and MEGAN.

2.3.2 Vegetation variables for the BVOC modules

Leaf area index: Measurements of the number of leaves in the canopy are required for ecosystem studies such as this one. This metric is often defined as the leaf area index (LAI), which is the one-sided leaf area in the canopy per unit surface area of the ground ($m^2\ m^{-2}$) (Jordan, 1969). In DGVMs, including LPJ-GUESS, this is a standard output variable.

135 **Foliar density:** The foliar density D (g dry matter m^{-2}), sometimes referred to as dry matter (DM), can be derived directly from the LAI as follows (Guenther et al., 1995):

$$D = \text{LAI} \cdot S_{LW} \tag{1}$$

where S_{LW} is an average specific leaf weight (g m^{-2}) and is given for each ecosytem (or PFT) based on Box (1981).

Leaf area density distribution: The LAD is a metric describing the leaf area in a cubic volume within the canopy (m^2
140 m^{-3}). The original ONEMIS configuration employs an *expert-driven* offline dataset from a DDIM (dry deposition inferential model) (Weiss and Norman, 1985) to characterise the LAD distribution for three types of vegetation: (i) agricultural crops, (ii) deciduous forests, and (iii) coniferous forests. The twelve PFTs in LPJ-GUESS were classified into these three groups, with grass PFTs included in the "agricultural crop" category. The LAD distribution for each of the vegetation types is divided into four equal layers; LAD 1 (top layer), LAD 2, LAD 3 and, LAD 4 (bottom layer).
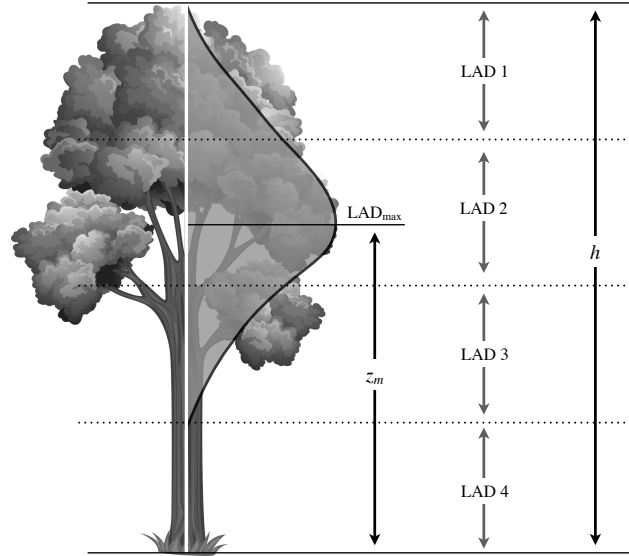
Figure 2. Graphical representation of the LAD distribution.

These values are then used in ONEMIS for calculating the photosynthetic active radiation (PAR) within the canopy and subsequent BVOC emission estimates. For the LPJ-GUESS output, we employ a parametrisation, derived by Lalic et al. (2013), to compute similar LADs at four canopy levels in the interface submodel between EMAC and LPJ-GUESS. The empirical relation describes the LAD at a height z (in m) as a function of the maximum LAD (LAD_{max}), the canopy height h, and the height z_m corresponding to LAD_{max} (see Fig. 2).

$$\text{LAD}(z) = \text{LAD}_{max} \left(\frac{h - z_m}{h - z} \right)^n \exp\left[n \left(1 - \frac{h - z_m}{h - z} \right) \right] \qquad \begin{array}{ll} n = 6 & \text{for} \quad 0 \leq z < z_m \\ n = \frac{1}{2} & \text{for} \quad z_m \leq z \leq h \end{array} \qquad (2)$$

We first extract the ratio between the canopy height (h) and the height corresponding to LAD_{max} (i.e. h/z_m) for each vegetation class using LAD canopy profiles from the DDIM. The dataset has 21 layers, and the layer where LAD_{max} occurs (z_m) is utilised to compute h/z_m as follows:

(i) Agricultural crops : $z_m/h = 12/21 \approx 0.57$

(ii) Deciduous forests : $z_m/h = 15/21 \approx 0.71$

(iii) Coniferous forests : $z_m/h = 17/21 \approx 0.81$

LAD_{max} for each PFT is calculated from the corresponding LAI, h, and h/z_m, given the relation:

$$\text{LAI} = \int\limits_{0}^{h} \text{LAD} = \int\limits_{0}^{h} \text{LAD}_{max} \left(\frac{h - z_m}{h - z} \right)^n \exp\left[n \left(1 - \frac{h - z_m}{h - z} \right) \right] dz \qquad (3)$$

After a numerical value for LAD_{max} for each PFT is computed, the LAD at four canopy layers is calculated via Eq. 2 by integrating over four equal layers within the total canopy height h. In all setups used in the study, h was assumed to be 25 m. This gives

$$
\begin{aligned}
\text{LAD}(1) &= \int\limits_{0.75h}^{h} \text{LAD}_{max} \left(\frac{h - z_m}{h - z} \right)^n \exp\left[n \left(1 - \frac{h - z_m}{h - z} \right) \right] dz, \\[2mm]
\text{LAD}(2) &= \int\limits_{0.5h}^{0.75h} \text{LAD}_{max} \left(\frac{h - z_m}{h - z} \right)^n \exp\left[n \left(1 - \frac{h - z_m}{h - z} \right) \right] dz, \\[2mm]
\text{LAD}(3) &= \int\limits_{0.25h}^{0.5h} \text{LAD}_{max} \left(\frac{h - z_m}{h - z} \right)^n \exp\left[n \left(1 - \frac{h - z_m}{h - z} \right) \right] dz, \\[2mm]
\text{LAD}(4) &= \int\limits_{0}^{0.25h} \text{LAD}_{max} \left(\frac{h - z_m}{h - z} \right)^n \exp\left[n \left(1 - \frac{h - z_m}{h - z} \right) \right] dz.
\end{aligned}
\qquad (4)
$$

where z_m is a fraction of h based on the PFT's vegetation class i, ii, or iii. Fig. 3 compares the LAD distribution from DDIM point data, used in the previous setup, with the new parametrisation described in Eq. 2.

Vegetation class coverage: The vegetation coverage refers to the fraction of land area occupied by certain PFTs in one grid cell. This variable is used in MEGAN to adjust emission rates from different vegetation classes. This variable is already calculated in LPJ-GUESS for each of the twelve PFTs.
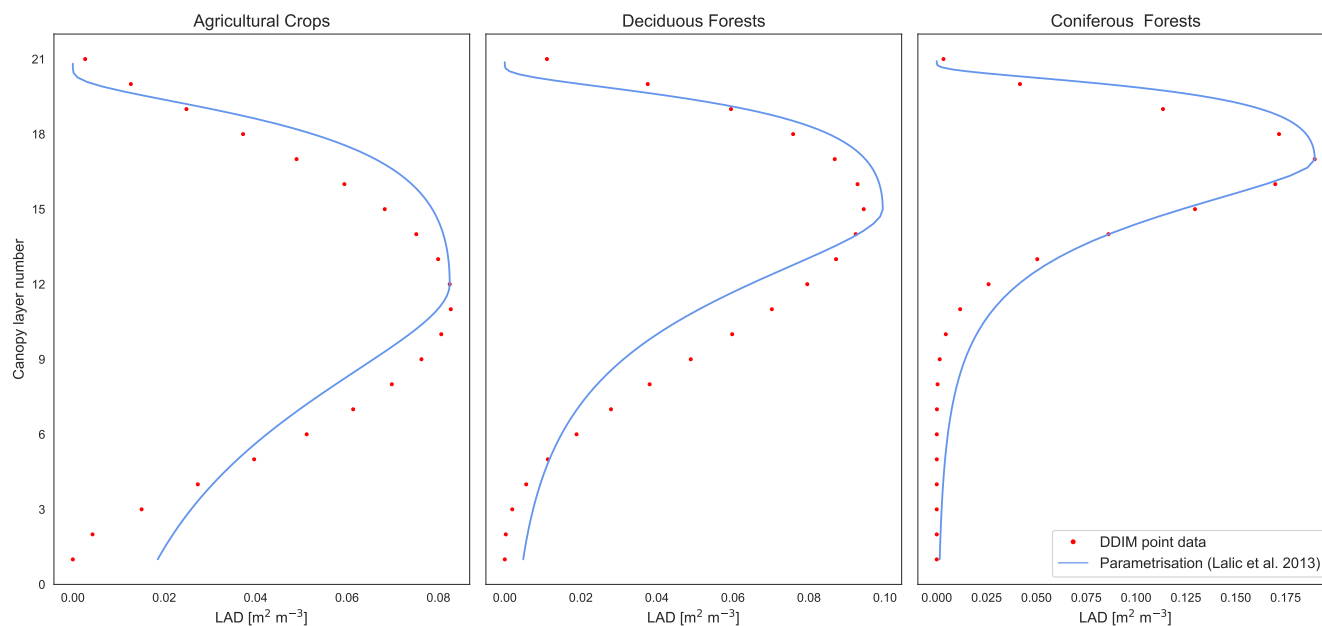
Figure 3. LAD distribution for a 21-layer canopy using DDIM point data versus the continuous distribution from the employed parametrisation for the three vegetation classes.

3   Results and discussion

3.1   Vegetation characteristics as input to the emission routines

170   In this section, we discuss the LPJ-GUESS state variables needed as input for the BVOC routines. We compare the LPJ-GUESS output with the corresponding offline climatological datasets used for BVOC emission estimates in ONEMIS and MEGAN in the original model configuration i.e. with prescribed vegetation boundary conditions. The simulation results presented are from a 10-year temporal average with a 500-year offline spin-up phase at a horizontal resolution of T63 (approximately $1.9° \times$ $1.9°$ at the Equator). The simulations are climatological, meaning that the same boundary conditions are used for each year of

175   the simulation.
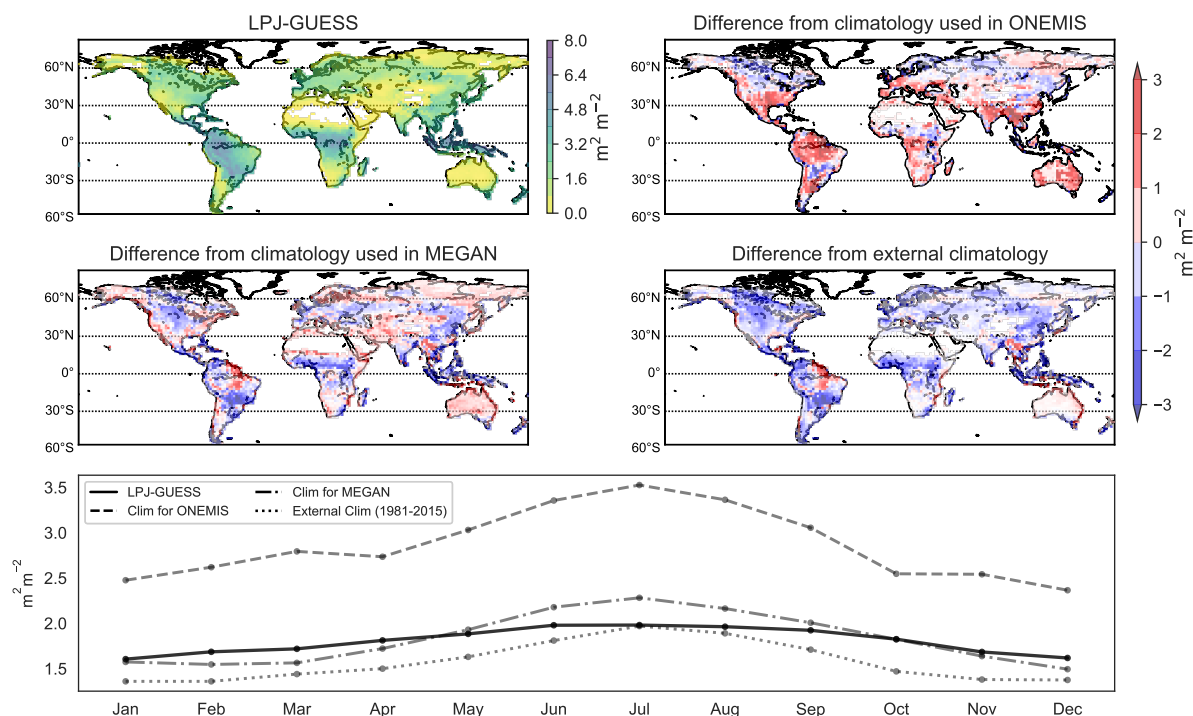
Figure 4. Geographic distribution of the LAI from LPJ-GUESS (top left panel) as well as difference plots (LAI from LPJ-GUESS minus climatological LAI) from climatology used in the standard configuration in ONEMIS and MEGAN. An additional LAI climatology from 1981 to 2015 is also included. The bottom panel compares the global monthly averages of all datasets.

The spatial distribution of the LAI from LPJ-GUESS - shown in Fig. 4 - indicates elevated LAI of more than 6 m^2 m^{-2} in the tropical rain forests of the Amazon, Congo and South East Asia. The LPJ-GUESS output is also compared with (1) the LAI prescribed for ONEMIS; (2) the LAI prescribed for MEGAN; and (3) an external climatology dataset (1981-2015) of the global monthly mean LAI (Mao and Yan, 2019). The difference plots indicate that the climatological LAI used in ONEMIS is

180 generally higher across the globe while the LAI climatology used in MEGAN is generally lower, especially over the tropics. The external climatological dataset prescribe lower LAI compared to the LAI from LPJ-GUESS, potentially due to the fact the LPJ-GUESS assumes a natural vegetation. The lower panel in Fig. 4 shows the global monthly mean LAI from all datasets. LPJ-GUESS exhibits a difference of 0.38 m^2 m^{-2} between the lowest and highest month, while the variability is 1.16 m^2 m^{-2} for the ONEMIS dataset, 0.79 m^2 m^{-2} for the MEGAN dataset, and 0.61 m^2 m^{-2} for the external LAI dataset. This means

185 that the LAI variability from LPJ-GUESS is comparable to the LAI climatology used in MEGAN as well as the external LAI dataset, however, it significantly differs from the variability in the climatological LAI used in ONEMIS. The foliar density is by definition a function of the LAI (Eq. 1), and hence shows a similar spatial and temporal distribution.

Fig. 5 displays the LAD distribution from LPJ-GUESS and climatology. The climatological LAD is from a DDIM model and was used in the original setup in ONEMIS to calculate BVOC emission estimates whereas the LAD from LPJ-GUESS is

190 derived from the parametrisation discussed in section 2.3.2. In LPJ-GUESS, a total canopy height of 25 m was assumed and thus the four LAD layers are classified as follows: bottom canopy layer (LAD 4) represents the LAD between 0-6.25 m, LAD 3 from 6.25-12.50 m, LAD 2 from 12.50-18.75 m, and the top canopy layer (LAD 1) from 18.75-25 m. Values of the four LAD layers (i.e. total canopy) adds up to one. The four panels on the right in Fig. 5 shows the geographic distribution of LAD from climatology (DDIM model). Even though, the LAD canopy distribution from climatology data is over-simplified it clearly

195 indicates higher densities in leaf area in the upper most layers of the canopy over the forested regions found in the tropics and boreal forests in the northern hemisphere. The panels on the left show the LAD geographic distribution from LPJ-GUESS at T63. Here we see a better-resolved geographic distribution for the LAD at the different canopy layers. Increased LAD in the bottom layer (i.e. all vegetation below 6.25 m) clearly highlight grasslands and desert areas. The layers above show how the LAD changes at the upper sections of the canopy, with plant biomass higher than 20 meters mostly found in tropical forest

200 areas.
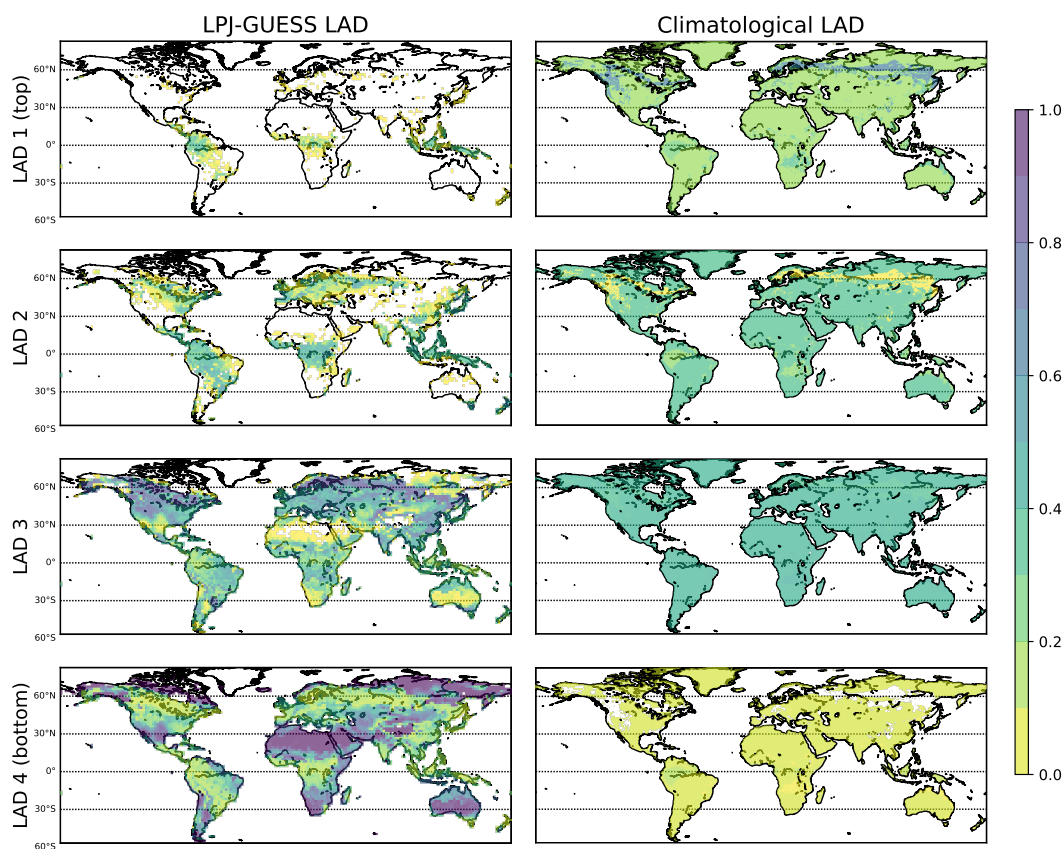
Figure 5. LAD fraction at four canopy layers from LPJ-GUESS (left panels) and from the DDIM model (right panels).
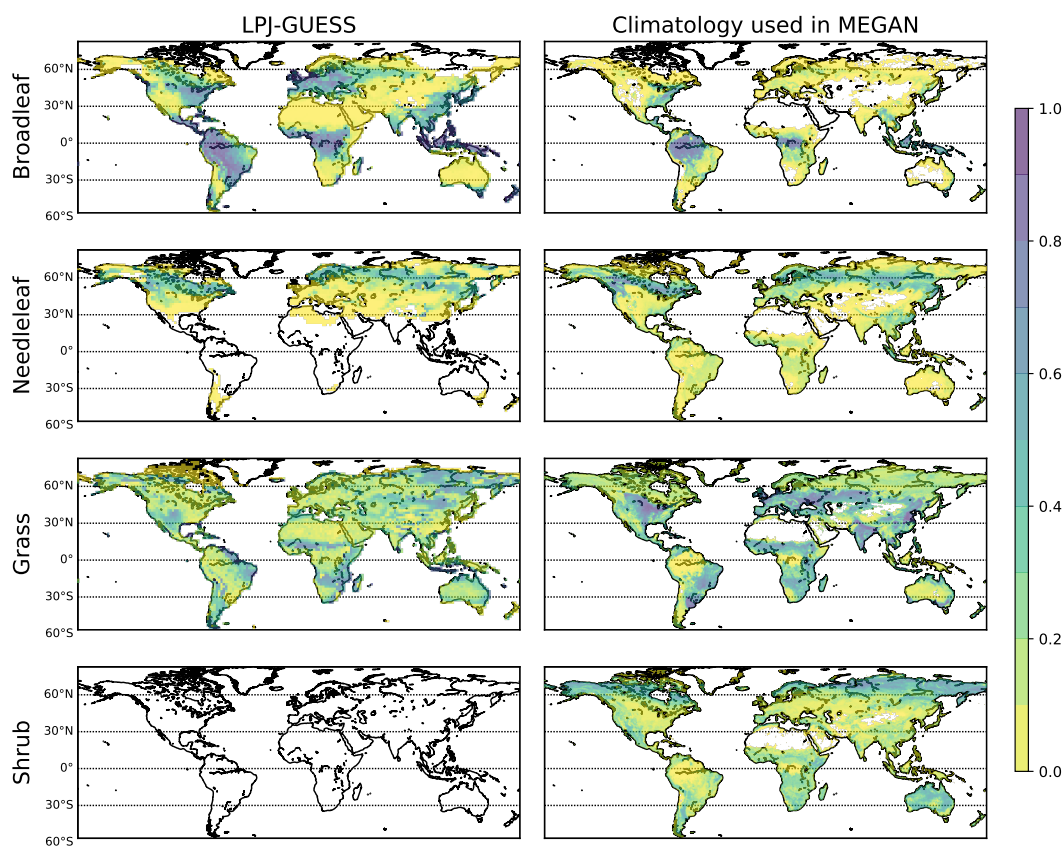
Figure 6. Fractional coverage for different vegetation classes used as inputs in MEGAN.

Fig. 6 shows the fractional coverage required as input for MEGAN. The twelve PFTs from LPJ-GUESS were classified in the four vegetation classes: broadleaf, needleleaf, grass and shrub, and are compared to the corresponding climatology vegetation fraction. A limitation of this setup is that LPJ-GUESS does not include shrub PFTs and the needleleaf representation in mixed forests is also missing.

205 **3.2 Global isoprene and monoterpene emissions**

Isoprene emissions

Fig. 7 presents global isoprene emissions from ONEMIS and MEGAN with dynamic vegetation states from LPJ-GUESS as well as offline climatological vegetation inputs, at spatial resolution T63. The simulations were conducted with fixed prescribed CO_2 of 348 ppmv (in both the radiation and vegetation schemes), and were performed using exactly the same code and 210 parameter settings, only changing the input *vegetation* parameters for the BVOC modules. All simulations are climatological and the results presented here are from 10 ensemble years.
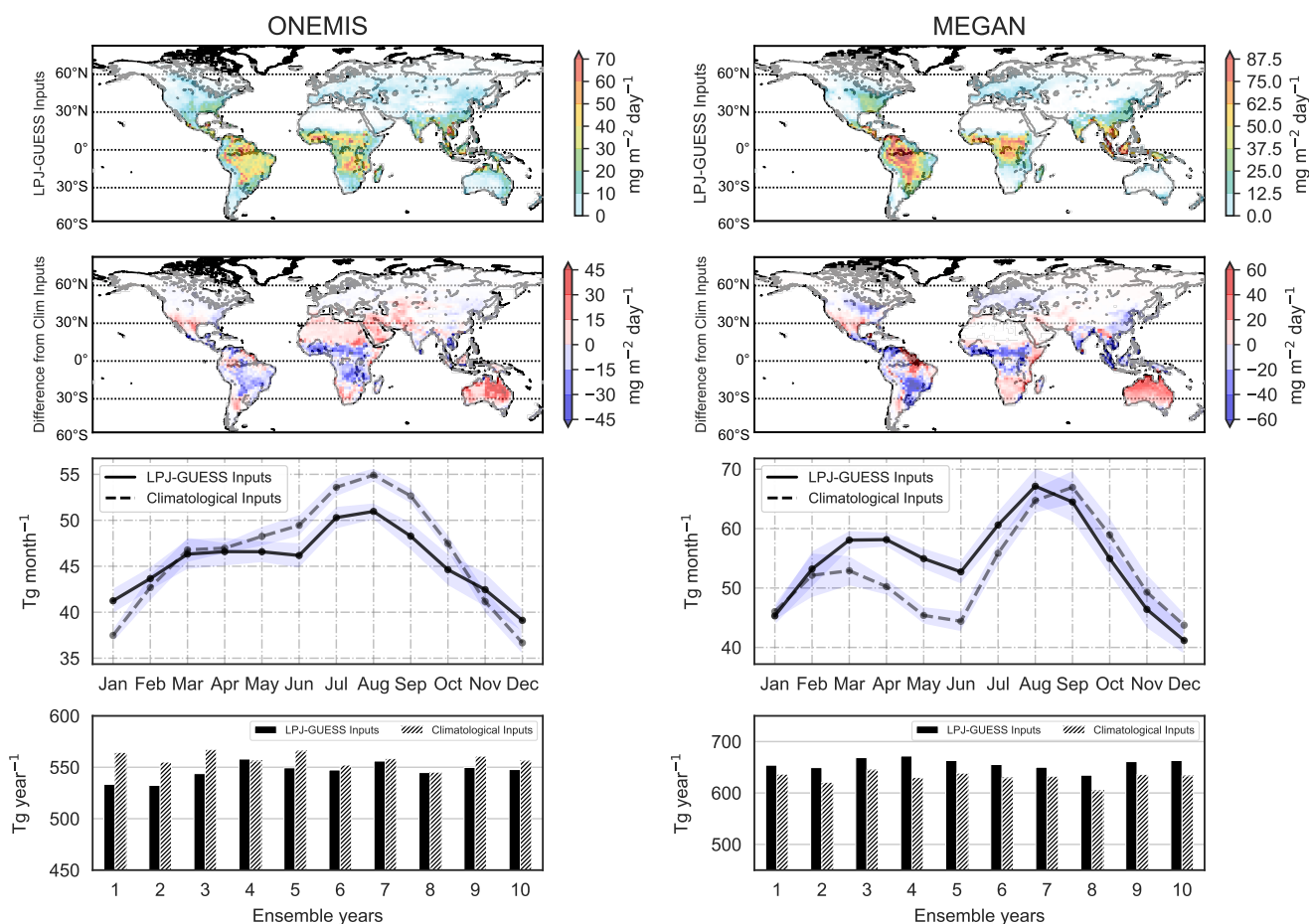
Figure 7. Spatial distribution of monthly isoprene emissions (mg m^{-2} day^{-1}) averaged over 10 ensemble years with LPJ-GUESS vegetation inputs from ONEMIS and MEGAN (top panels). The panels in the second row show the difference in the emission flux compared to climatological inputs. Temporal profile of isoprene global totals (Tg month^{-1}) (third row panels). Global annual totals (Tg yr^{-1}) for 10 years (bottom panels).

The top panels in Fig. 7 show the geographic distribution of isoprene emission rates (in mg m^{-2} day^{-1}) using LPJ-GUESS vegetation inputs. Elevated isoprene emissions can be seen over South America, Central Africa, Southeast Asia, and North Australia mostly corresponding to high vegetation densities in tropical rain forests. MEGAN emissions are significantly higher with up to 90 mg m^{-2} day^{-1} over the Amazon forest. The plots in the second row of Fig. 7 compare our new emissions using LPJ-GUESS vegetation with emissions using climatological inputs. For both ONEMIS and MEGAN, emissions with dynamic vegetation are lower over tropical areas but higher over Australia compared to emissions using climatology. With climatologcial vegetation inputs, ONEMIS exhibits low but significant emissions over deserted regions, particularly the Saharan desert in North Africa, resulting from the poor representation of the LAD distribution in the original model setup. This is not the case when using dynamic vegetation inputs.

The middle panel in Fig. 7 depicts the temporal profile of global monthly emission totals, whereas the bottom panel shows the inter-annual variability of isoprene annual global totals. Over the ten ensemble years, isoprene emissions appear to peak in the boreal summer months and decrease substantially in the boreal winter months. The stronger seasonal variability in ONEMIS fluxes with climatological inputs results from a corresponding stronger variability in the LAI inputs (see Fig. 4).

225 We are confident that the variability in isoprene emissions in the new model setup is more realistic considering that the LAI from LPJ-GUESS is more comparable to recent climatologies. MEGAN includes a leaf age factor which accounts for reduced emissions for young and old leaves based on observed LAI change. This algorithm in MEGAN delays the peak in isoprene emissions. The climatological LAI used for MEGAN varies more significantly from May to June to July compared to LAI from LPJ-GUESS (see Fig. 4). The higher monthly LAI change amplifies the effects of the leaf age factor which explains the

230 stronger reduction in emissions with climatological LAI.

Over the 10-year simulation period considered, the global annual total isoprene fluxes from ONEMIS were found to be 546 Tg yr^{-1} (SD = 8 Tg yr^{-1}) with dynamic vegetation and, 558 Tg yr^{-1} (SD = 7 Tg yr^{-1}) with climatological inputs. MEGAN estimated 657 Tg yr^{-1} (SD = 11 Tg yr^{-1}) with dynamic vegetation and 631 Tg yr^{-1} (SD = 11 Tg yr^{-1}) with climatological vegetation inputs. The higher standard deviation when using LPJ-GUESS inputs indicates higher inter-annual variability in

235 isoprene fluxes. While the year-to-year variability with climatological inputs is only determined by surface temperature and short wave radiation (see Fig.1), the cross-annual variability with LPJ-GUESS inputs is also influenced by inter-annual changes in vegetation states estimated in LPJ-GUESS. Jöckel et al. (2016) reported isoprene annual emissions of 488-624 Tg using ONEMIS, while other studies estimated fluxes of 642 Tg yr^{-1} (Shim et al., 2005) using 73 prescribed vegetation types, 571 Tg yr^{-1} (Guenther et al., 2006) using inventories and Olsen ecoregions land covers, 467 Tg yr^{-1} (Arneth et al., 2007a) using 10

240 PFTs from LPJ-GUESS, and more recently, 594 Tg yr^{-1} using 16 PFTs (Sindelarova et al., 2014). It is important to note that no scaling factors were applied in our simulations, even though global annual totals from models are often scaled in atmospheric chemistry studies. For example, Pozzer et al. (2022) estimated 464 Tg of isoprene in the year 2010 using a MEGAN-EMAC setup but employed a global scaling factor of 0.6. This means that the original emissions are similar to our MEGAN fluxes with climatological inputs.

245 **Monoterpene emissions**

Results for monoterpene emissions are shown in Fig. 8. The top panels present the spatial distribution of monoterpene emission rates in mg m^{-2} day^{-1} from ONEMIS and MEGAN, while the the panels in the second row show difference in the emission fluxes using climatological vegetation inputs. Here we also observe elevated emission rates over tropical rain forest areas, with ONEMIS prescribing much higher emission rates compared to MEGAN. The difference plots indicate that

250 monoterpene emissions from ONEMIS are significantly higher with climatological inputs except for some areas over Southern Brazil and Central Africa. Similarly, MEGAN generally prescribes higher emissions with climatological vegetation inputs compared to LPJ-GUESS inputs.ONEMIS emission over deserted regions in north Africa and central Australia appear to be better resolved with dynamic vegetation, since ONEMIS with climatology inputs still attributes significant emission rates over such regions where vegetation is absent.
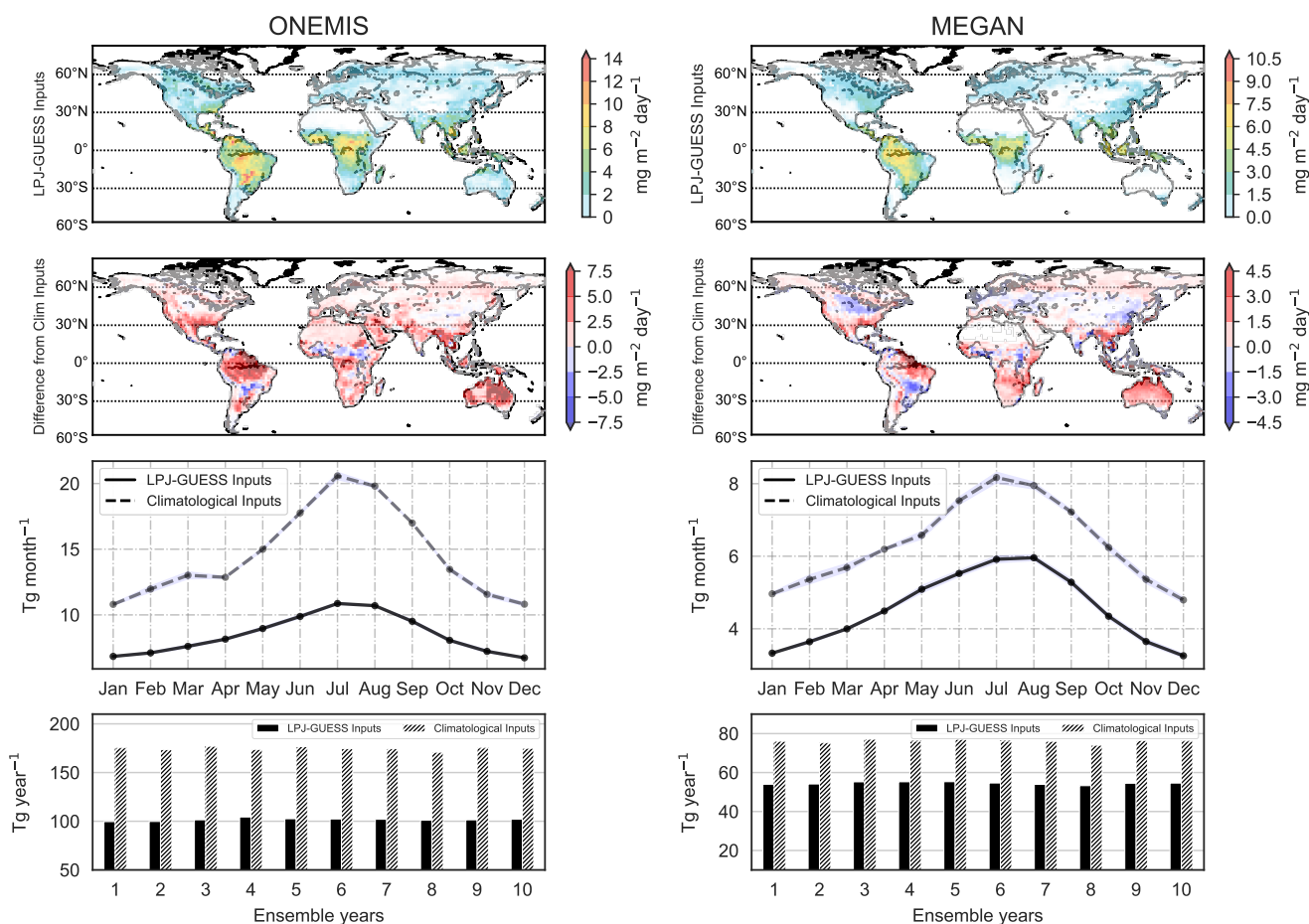
Figure 8. Same as Fig. 7 but for monoterpenes.

Monoterpene emission fluxes in ONEMIS only depend on the foliar density profile (i.e. DM × LAD) and surface temperature. The high dependence of monoterpene emissions on the foliar density explains both the lower seasonal variability as well as the lower yearly fluxes compared to emissions with climatological vegetation inputs. The lower magnitudes in monoterpene fluxes from MEGAN with dynamic vegetation results from the lack of representation of shrubs and needleleaf PFT's in LPJ-GUESS, where such species are know to be strong emitters of monoterpenes. The seasonal variation, however, is satisfying, and the fact that the fractional coverages computed in LPJ-GUESS are dynamic (updating on a yearly time-step) makes this setup suitable for studying long-term variations in emissions. Annual totals from ONEMIS were found to be 102 Tg yr^{-1} (SD = 1 Tg yr^{-1}) with dynamic vegetation inputs and 175 Tg yr^{-1} (SD = 2 Tg yr^{-1}) with climatological inputs. MEGAN prescribes 54 Tg yr^{-1} (SD = 0.7 Tg yr^{-1}) and 76 Tg yr^{-1} (SD = 0.9 Tg yr^{-1}) with dynamic and climatological inputs respectively. Guenther et al. (2012) gives global annual monoterpene emission of 157 Tg, while Sindelarova et al. (2014) reported annual

265 total emissions of monoterpenes to range between 89 and 102 Tg yr^{-1} over a 30-year simulation period. Arneth et al. (2007a) reported 36 Tg yr^{-1}.

LPJ-GUESS BVOC emission routine

LPJ-GUESS includes a built-in BVOC emission module for the calculation of isoprene and monoterpene emission fluxes. The submodel combines the process-based leaf level emission model (Niinemets et al., 2002, 1999) with the LPJ-GUESS

270 vegetation model for isoprene (Arneth et al., 2007b) and monoterpene (Schurgers et al., 2009) emissions. The algorithm computes BVOC production based on photosynthetic electron flux, emission factors, temperature, seasonality and also includes a CO_2 inhibition factor on leaf production of isoprene and monoterpenes relative to the ~ 370 ppmv [CO_2] in the year 2000 (Hantson et al., 2017). The algorithm also needs the daily temperature range (DTR) for the calculation of BVOC emission rates, which is typically derived from climatological datasets. In this study, we extend the coupling between EMAC and LPJ-

275 GUESS by computing the DTR in EMAC (defined as the difference between the maximum and minimum daily temperature) and feeding it to LPJ-GUESS on a daily time-step.
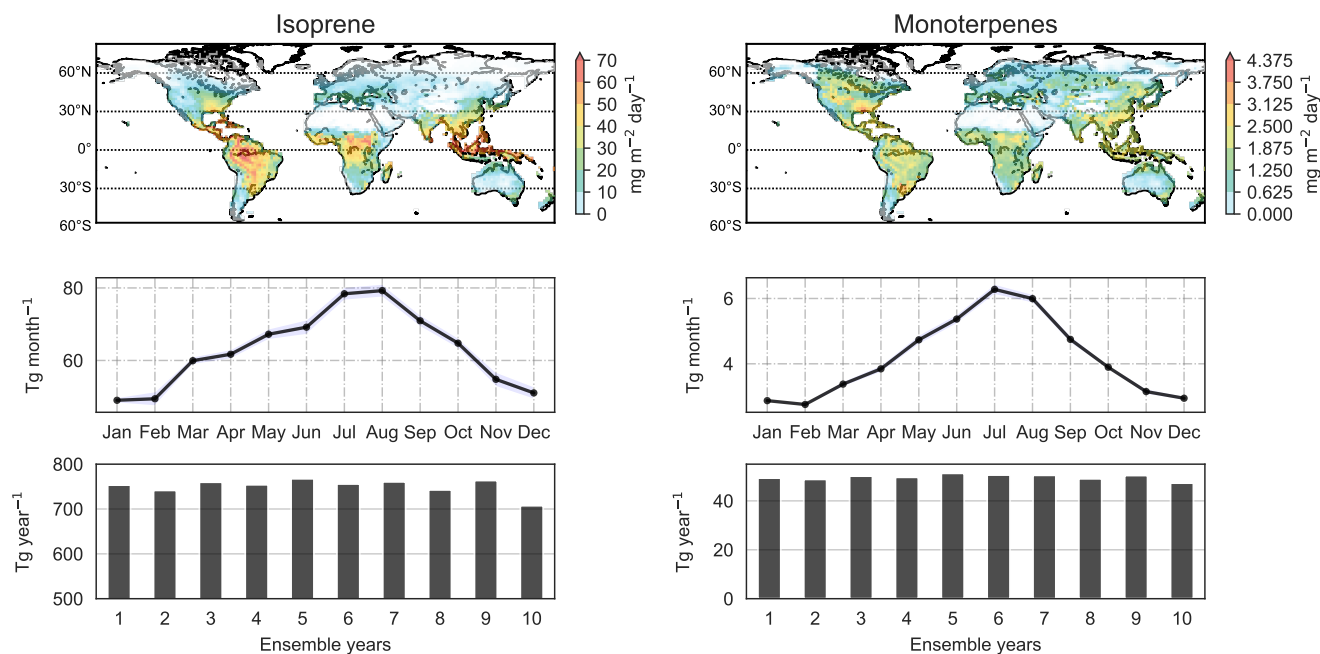
Figure 9. Spatial distribution of isoprene (top left) and monoterpene (top right) emissions from LPJ-GUESS at spatial resolution of T63. Isoprene and monoterpene mean monthly totals in Tg month^{-1} (mid panels) and annual totals in Tg yr^{-1} for 10 ensemble years (bottom panels).

The top panels in Fig. 9 show the spatial distribution of isoprene and monoterpene emission rates from the LPJ-GUESS BVOC emissions routine. Isoprene monthly emissions range from 49.0 Tg in January to 79.3 Tg in July, whereas monthly monoterpene emissions range from 2.8 Tg in February to 6.3 Tg in July. The lower panels show yearly isoprene and monoter-

280 pene totals from LPJ-GUESS for 10 ensemble years. The mean isoprene annual emission flux is 750 Tg yr^{-1} (SD = 17 Tg yr^{-1}) while for monoterpenes it is 50 Tg yr^{-1} (SD = 1 Tg yr^{-1}). Hantson et al. (2017) evaluated the LPJ-GUESS algorithm and reported global annual isoprene and monoterpene emissions of 454 Tg yr^{-1} and 34 Tg yr^{-1}, respectively.

3.3 Emission sensitivity to double CO$_2$ scenarios

In this section, we investigate the variability of global isoprene and monterpene emission estimates in doubling CO$_2$ scenarios.
285 In particular, we evaluate CO$_2$-fertilisation and temperature effects by prescribing different CO$_2$ values in the radiation and vegetation schemes in the coupled model setup. The submodel RAD in EMAC (Dietmüller et al., 2016) simulates the radiative transfer in the atmosphere accounting for the effects of the shortwave and long-wave radiation fluxes from radiatively active trace gases. CO$_2$ has the largest radiative influence in the longwave range of the spectrum, resulting in radiative forcings leading to stratospheric cooling and tropospheric warming. The CO$_2$ value prescribed in RAD mainly dictates surface temperatures
290 resulting from the greenhouse effect, while CO$_2$ in the vegetation scheme (i.e. in LPJ-GUESS) determines the carbon available for photosynthesis and hence accounts for CO$_2$-fertilisation effects. Four experiments were conducted to explore the impact of doubling CO$_2$ scenarios (both in the radiation and vegetation scheme) on isoprene and monoterpene emission rates. Ref is the reference simulation, with 348 ppmv [CO$_2$] in both schemes. Table 1 lists the CO$_2$ levels prescribed for each study.

Study	CO$_2$ in radiation scheme	CO$_2$ in vegetation scheme
Ref	348 ppmv	348 ppmv
Bio×2	348 ppmv	696 ppmv
Atm×2	696 ppmv	348 ppmv
Both×2	696 ppmv	696 ppmv

Table 1. Prescribed CO$_2$ in the radiation and vegetation schemes for different studies.

Climatological monthly average sea surface temperature (SST) and sea ice content (SIC) from the AMIP database from
295 1987 to 2006 are used for Ref and Bio×2, with a prescribed CO$_2$ of 348 ppmv in the radiation scheme. However, with 696 ppmv [CO$_2$] in the radiation scheme (i.e. in Atm×2 and Both×2), oceanic boundary conditions are prescribed using external data corresponding to SST and SIC at 696 ppmv in order to preserve radiative equilibrium. This data is aquired from a coupled atmosphere–ocean general circulation model (increased/decreased SSTs/SICs) performed under identical climate circumstances (696 ppmv [CO$_2$]) (same approach as in Rybka and Tost (2014)). The analyses shown are based on ten years of
300 data from 50-year simulations with constant boundary conditions and a 500-year offline spin-up phase.

Surface temperature in 2×CO2 in the radiation scheme

We study temperature sensitivity of BVOC emissions by doubling the prescribed CO$_2$ value in the radiation scheme. This results in a consistent global increase in the surface temperature of up to 4°C. See Appendix.

Vegetation response to 2×CO2 scenarios

305 The LAI can be used as a marker for vegetation activity. Fig. 10 displays LAI estimates from LPJ-GUESS for the reference study, Ref, and also shows how LAI values in the other studies compare to it. Bio×2 indicates consistent increased vegetation activity when doubling the prescribed CO_2 in the vegetation scheme. This is the CO_2-fertilisation effect. In Atm×2 we see a decline in the LAI, resulting from warmer temperatures and higher losses of soil moisture (reported elsewhere e.g. Dermody et al., 2007; Sun et al., 2015). In Both×2, with 696 ppmv $[CO_2]$ in both the vegetation and radiation scheme, we find an overall

310 rise in LAI compared to the Ref simulation, except for some places in North America, Western Brazil and Southern Europe.
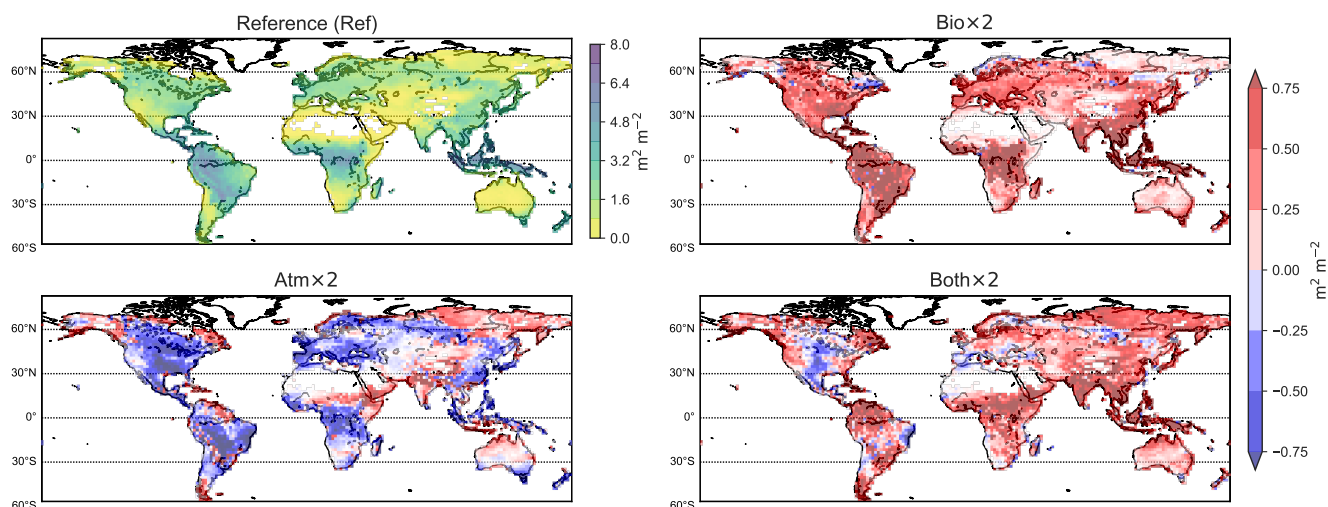
Figure 10. Geographic distribution of the LAI for the reference study (Ref) and difference plots for Bio×2, Atm×2, and Bio×2.

Global BVOC emissions

Fig. 11 shows the geographic distribution of isoprene emission rates (in mg m^{-2} day^{-1}) averaged over 10 ensemble years for Ref and difference plots for Bio×2, Atm×2, and Both×2 using climatological and dynamic vegetation inputs using ONEMIS and MEGAN. The bottom panels show the averaged emission rates for distinct latitude bands (0°- 30°S, 0°- 30°N, 30°N -

315 60°N) for all studies during the same time period, from left to right. We see that in Bio×2, isoprene emissions increased only when using dynamic vegetation inputs in both ONEMIS and MEGAN. In contrast to the prescribed climatological vegetation data, the LPJ-GUESS coupled setup is sensitive to increased CO_2 which subsequently leads to higher emissions. In Atm×2 we see temperature effects on isoprene emissions while in Both×2 we see the the combined effects of CO_2 fertilisation and temperature. The bar plots reveal that in places where most of the global emissions occur, i.e. between 0°- 30°S, the emission

320 was the highest in Both×2 with and average emission rate of 27.3 mg m^{-2} day^{-1} (ONEMIS) and 44.3 mg m^{-2} day^{-1} (MEGAN). Given the empirical nature of ONEMIS and MEGAN, both setups give a consistent increase in isoprene fluxes in elevated temperatures (Atm×2). However, with LPJ-GUESS inputs we can see differences in the emission fluxes also resulting from vegetation dynamics e.g. decrease in fluxes from lower vegetation activity caused by water stresses. This indeed shows the advantage of having BVOC fluxes derived from dynamic vegetation states rather then prescribed boundary conditions.
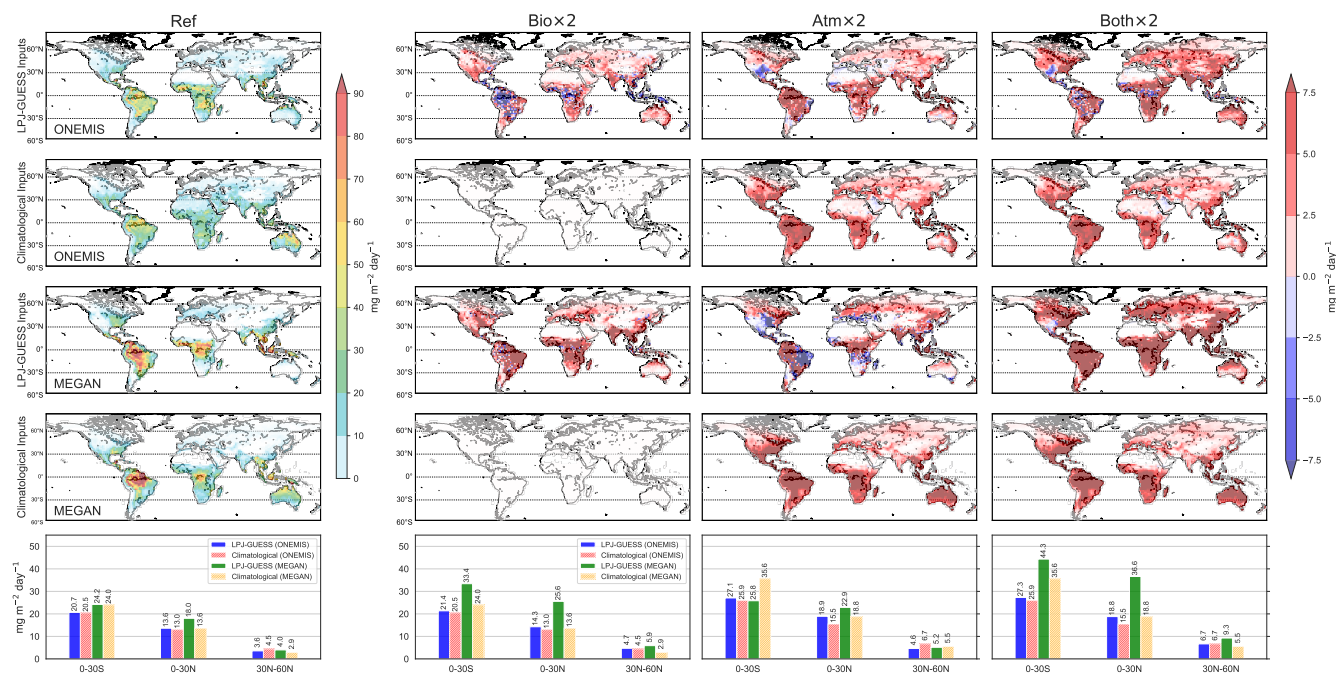
Figure 11. Global isoprene emission estimates from ONEMIS and MEGAN with LPJ-GUESS vegetation inputs and climatological inputs for the reference study (Ref) (left panels), as well as difference plots for Bio×2, Atm×2, and Both×2 (right panels). The panles in the bottom row displays emissions flux averages (in mg m^{-2} day^{-1}) over the latitude bands 0°- 30°S, 0°- 30°N, 30°N - 60°N.

325   In Fig. 12 we see similar behaviour in monoterpene emission estimates for all studies. Monoterpene emission rates increase in Bio×2 only for scenarios with dynamic vegetation due to CO_2-fertilisation. We detect a worldwide increase in monoterpene fluxes in Atm×2 with higher surface temperatures using climatological inputs, however with dynamic vegetation inputs, we see a less substantial rise as well as a drop in fluxes in certain regions. In Both×2 we see an increase in monoterpene emission rates as a result of the combined influence of temperature and CO_2-fertilisation.
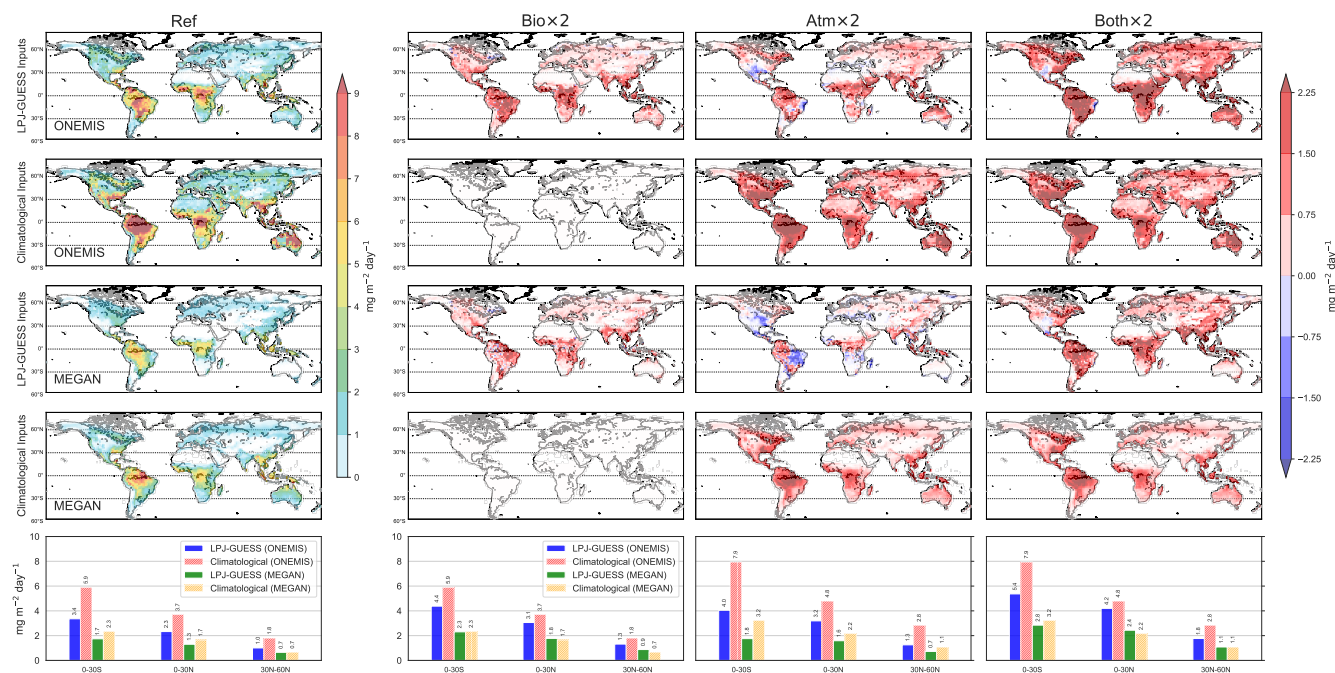
Figure 12. Same as Fig. 11 but for monoterpenes.

4 Conclusions

330   The complexities between plants and the atmosphere need better understanding and is an exciting research frontier. The development of ESMs will allow a more intricate analysis of a fully coupled and dynamic system addressing many complex biosphere-atmosphere interactions governed by BVOC emissions and thus shedding more light on the significance of such processes on global climate change and air-quality. In this work, further development was made towards producing a new

335   atmospheric chemistry-enabled ESM by integrating an atmospheric chemistry-enabled atmosphere–ocean general circulation models (AOGCM) with a DGVM. This work also explores, for the first time, partial bi-directional interactions between the two modelling systems via BVOC emissions, building on recent technical developments (Forrest et al., 2020) that enabled for one-way coupled simulations (in which climate information generated by EMAC is used to force LPJ- GUESS, but no land surface information is relayed back to EMAC).

340   Results show that the LAI and subsequent foliar density estimations from LPJ-GUESS are comparable to climatological datasets used as boundary layer conditions in the MEGAN-EMAC stand-alone configuration, as well as an external LAI climatolgy from 1981 to 2015. The LAI employed in the original ONEMIS-EMAC setup, on the other hand, differs significantly in terms of magnitude and monthly variability. Such differences in the LAI inputs led to lower and less-varying isoprene and monoterpene emissions using the coupled model setup compared to the stand-alone configuration. The representation of the

345 LAD distribution from the new parametrisation employed in LPJ-GUESS also gives a more realistic LAD profile compared to the over-simplified datasets used by the standard ONEMIS setup.

The new MEGAN-LPJ-GUESS configuration yielded satisfactory isoprene fluxes as well, however they are marginally higher than emissions using climatological inputs due to MEGAN's leaf-ageing effects, which are less prominent when the difference in the prescribed monthly LAI is smaller. Monoterpene emissions from the MEGAN-LPJ-GUESS setup also gives 350 meaningful monthly variability with lower magnitudes compared to emissions with climatological inputs given that the new setup lacks emission contributions from shrubs and needleleaf PFTs in mixed forests.

Global isoprene emission estimates from the coupled model configuration, using both ONEMIS and MEGAN, give realistic global variability, corresponding to emissions with prescribed vegetation boundary conditions. The emission magnitudes are also comparable to modelled fluxes found in literature and can be adjusted accordingly with global scaling factors for specific 355 atmospheric chemistry studies.

CO_2 sensitivity studies also suggest that both isoprene and monoterpene emissions rise with warmer climatic scenarios ($2\times CO2$ in the radiation scheme), however only the coupled configuration showed reduced emissions in some locations. The lower emissions result from the vegetation response to water stresses in a warmer climate. The new coupled model also allows for sensitivity studies for CO_2-fertilisation effects and indicate an increase in both isoprene and monoterpene emission rates in 360 $2\times CO2$ scenarios in the vegetation scheme due to enhanced vegetation activity in a CO_2 rich-atmosphere. This work provides evidence that the improved ESM, featuring dynamic vegetation, gives realistic BVOC emission estimates on a global scale based on dynamic vegetation states, enabling for further research into atmosphere-biosphere interactions and feedbacks with this model configuration.

Code availability. The Modular Earth Submodel System (MESSy) is continuously developed and applied by a consortium of institutions. 365 MESSy is licensed to all affiliates of institutions that are members of the MESSy Consortium, as is access to the source code. Institutions can become a member of the MESSy Consortium by signing the MESSy Memorandum of Understanding. More information can be found on the MESSy Consortium website (http://www.messy-interface.org, last access: 28 April 2022). LPJ-GUESS is used and developed globally, however development is overseen at Lund University's Department of Physical Geography and Ecosystem Science in Sweden. Model code can be made available to collaborators on entering into a collaboration agreement with the acceptance of certain conditions. Given that both 370 MESSy and and LPJ-GUESS are restricted, the code used here is archived with a restricted access DOI (10.5281/zenodo.6772205). The code will not be made publicly available and sharing will be made possible only by the approval of the authors.

Appendix A: Surface temperature at doubling CO_2 scenarios

Fig. A1 illustrates temperatures (°C) at the lower most vertical level in EMAC at spatial resolution T63 with prescribed CO_2 of 348 ppmv in the top panel and 696 ppmv in the middle panel. Plot comparing the reference simulations to the $2 \times CO2$

375   simulations is given in the lower panel of Fig. A1.
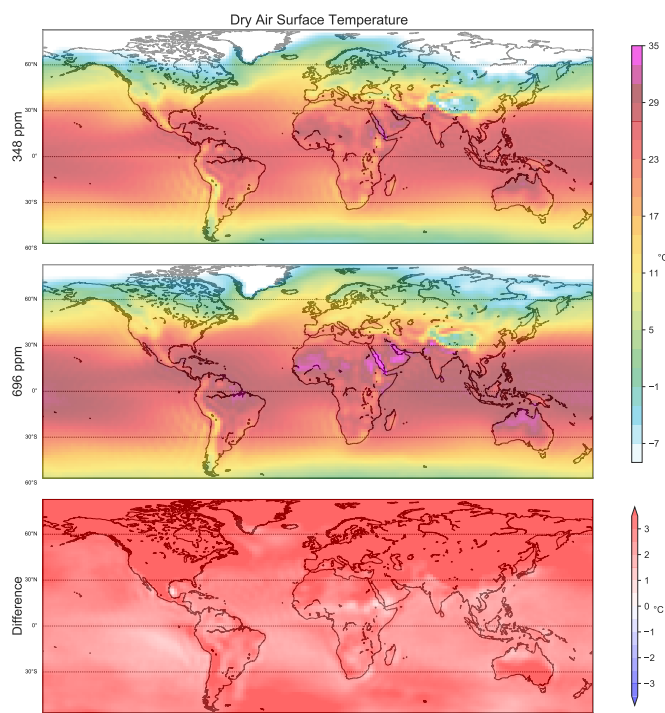
Figure A1. Surface temperature (°C) at spatial resolutions T63. In the radiation scheme, the first model configuration utilises a CO_2 volume mixing ratio of 348 ppm, whereas the second setup uses 696 ppm.

Author contributions. RV and HT performed the model coupling. RV performed the simulations and analysis. All authors contributed to the overall model development, scientifc analysis and wrting of the article.

Competing interests. HT acts as a topical editor for GMD. Apart from this, the authors declare that they have no conflict of interest.

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