

## Response to Reviewer 1:

Review of “CANOPS-GRB v1.0: a new Earth 1 system model for simulating the evolution of ocean-atmosphere chemistry over geologic timescales”

This paper describes a new version of the CANOPS Earth system model, CANOPS-GRB (Global Redox Budget). The main improvements of this model over previous versions include dynamic atmospheric O<sub>2</sub>, sedimentary reservoirs that evolve self-consistently with the atmosphere-ocean system, and a simplified global redox budget. Collectively, these improvements make the model suitable for investigating long timescale geochemical evolution, but with the benefits of resolving ocean transport and sediment diagenesis. Model outputs are validated against the observed modern Earth steady state, and example time-evolution calculations are also presented to illustrate feedbacks on multiple timescales.

I will preface this review by saying evaluating the validity of model assumptions in the abstract is challenging; the appropriateness of the chosen functions, parameterizations, and simplifications are strongly context dependent. However, I have no doubt that this manuscript and open-source model it describes will be incredibly valuable to the biogeochemistry community. The authors ought to be commended for such a thorough description of model features – this will be invaluable for trying to understand future applications of the model. I don't have any major criticisms – the primary limitations of the model are largely already identified and discussed in the manuscript. However, I do have several minor suggestions for improvement:

We are grateful to Reviewer 1 for the useful suggestions and critical comments, which have helped us to significantly improve our paper. As indicated in the responses that follow, we have taken all these comments and suggestions into account in the process of revision. We are pleased to say that the comments and suggestions have both greatly improved the quality of the manuscript.

The Introduction motivates the CANOPS-GRB model through a discussion of terrestrial exoplanets and exoplanet biosignatures such as oxygen and methane. Moreover, the abstract mentions exoplanet biogeochemistry as one potential application of CANOPS-GRB. It would be helpful if the discussion also had a paragraph or two on both the potential and limitations of the current model for terrestrial exoplanets. For example, by adopting the Claire et al. photochemical parameterization and by omitting a climate model, the ability to investigate O<sub>2</sub>-CH<sub>4</sub> biosignatures is limited to Earth-twins around a sun-like stars (see also line 351). Obviously, these parameterizations could be modified to accommodate different stellar types etc., but I think it is helpful to make the current limitations of the model clear to the exoplanet community.

This is an important point. We have added brief discussion to the revised Discussion section to clarify this more explicitly (Line 1681-1686): “The CANOPS-GRB model has the basic capability

to simulate the time evolution of the abundance of atmospheric biosignature gases (O<sub>2</sub> and CH<sub>4</sub>) on a wide range of timescales. While the biogeochemical model is based on process studies to the extent possible, many processes are derived from empirical calibrations to Earth-like planets around sun-like stars. Clearly, some of these parameterizations, such as the photochemical parameterization among O<sub>2</sub>-O<sub>3</sub>-CH<sub>4</sub> (Eq. 107), must be modified when applying the model to a range of habitable Earth-like exoplanets.”

Line 229-230: I appreciate that the omission of Fe<sup>2+</sup> is discussed later in the paper, but it might be worth mentioning this here (e.g. see discussion of iron species below) since upon first reading I immediately wondered why the iron budget was being ignored.

This is also an important point. We have added a brief sentence to make this point clear (Line 235-236) and appreciate this opportunity to clarify this point.

Line 291-292: Neglecting the inorganic carbon cycle is an interesting choice that merits more discussion. To be clear, I understand the reasoning behind this simplification – carbon cycle feedbacks are uncertain and introduce more unconstrained complexity to an already complex model. However, changes in inorganic carbon cycling have been proposed as drivers of atmospheric oxygenation (e.g. Williams et al. 2019; Nature Communications), and the <sup>13</sup>C record provides a useful sanity check on any proposed oxygenation story. Neglecting the inorganic carbon cycle also means climate feedbacks are absent, which prevents the model from exploring the coevolution of life and the environment in many deep time contexts. More discussion of these limitations would be helpful in preventing misapplications or misunderstandings of the model.

We agree with the reviewer. We have added explanations behind this simplification in the revised manuscript (Line 301-303): “Although we ignore the inorganic carbon cycle here for clarity and in order to simplify interpretation of model results, the full coupling of the inorganic carbon cycle within CANOPS-GRB is an important topic for future work.”

Line 329: 14.3 Tmol/yr total organic carbon burial feels a bit on the high side. For example, see the compilation of estimates in Table 1 of Kipp et al. (2020; Global Biogeochemical Cycles). Similarly, organic weathering in CANOPS-GRB is about double most literature estimates. The discrepancy is noted on line 1368, but I am left wondering why such a high organic carbon flux is required. Given isotope mass balance, a high organic burial flux would also imply quite a high total carbon burial flux, which potentially conflicts with other literature estimates.

The reviewer is correct to point out that our estimate of organic C burial is relatively high compared to previous estimates (though some key estimates that are similar to ours are left out of the Kipp et al. summary (e.g., Wallmann and Aloisi, 2012; Wallmann et al., 2012)). To clarify this point, we have added a brief discussion to the revised Discussion section (Line 1622-1624): “In our

reference run the estimated organic carbon burial and oxidative weathering fluxes are relatively high compared to some previous estimates, though there remains significant uncertainty in globally integrated organic carbon weathering and burial fluxes.”

Wallmann et al. (2012). The global inventory of methane hydrate in marine sediments: A theoretical approach. *Energies*, 5, 2449.

Wallmann, K. and Aloisi, G. (2012). The Global Carbon Cycle: Geological Processes. In *Fundamentals of Geobiology* (eds A.H. Knoll, D.E. Canfield and K.O. Konhauser).

Equation 2: This is another place where it might be worth mentioning that climate feedbacks on silicate weathering are an important omission that could have consequences for global redox.

We have added a brief sentence in Line 396-399: “Note that our representation of P weathering ignores the effect of climate (Eq. 2). In the current version of the model the rate of P weathering is treated as one of the model forcings. Although ignoring the climate feedback on P mobility makes interpretation of the model results more straightforward, the incorporation of a climate-sensitive crustal P cycle is an important avenue for future work.”

Line 458: Although this is stated later in the manuscript, this discussion of sulfur oxidation might be a good place to mention that the model in its current form is not suitable for modeling the Archean atmosphere-ocean system.

We added a brief sentence in the revised manuscript (Line 236-239): “All H<sub>2</sub>S and NH<sub>3</sub> degassing from the ocean to the atmosphere is assumed to be completely oxidized by O<sub>2</sub> to SO<sub>4</sub><sup>2-</sup> and NO<sub>3</sub><sup>-</sup> and returns to the ocean surface. These simplifications limit application of the model to very poorly oxygenated Earth system states ( $pO_2 < 10^{-5}$  PAL).”

Section 2.4.2: I appreciate that the vertical and horizontal transport parameterizations developed here are carefully validated later in the paper. However, one unanswered question I had is how well do the authors expect these parameterizations to work when they are applied to radically different surface climates in deep time (e.g. hothouse Earth, snowball Earth)? Some discussion of this might be helpful for readers considering applying the model to different times in Earth’s history (or indeed different exoplanet climates, continental configurations etc.)

Our water transport scheme is flexible enough to represent the general water circulation, such as thermohaline circulation and high-latitude convection, on any rotating planet, with a limited set of free parameters. However, the reviewer is correct to point out that the users need to be careful when applying the model to the ancient oceans or oceans on exoplanets. We added brief explanations of this point in the revised manuscript (Line 640-646): “Thermohaline circulation

and high-latitude convection are considered to be general physical modes on any rotating planet, and our simplified water transport scheme allows us to represent them with a limit number of free parameters. However, we emphasize that the water transport scheme explored here is designed to represent the modern ocean circulation on Earth. As a result, some of these parameterizations may need to be modified when applying to ancient oceans or oceans on exoplanets. Nevertheless, given our simple design, our water transport scheme is relatively flexible to modify the water circulations that are markedly different from the modern ocean.”

Equation 108 and 109: It’s possible I’m misunderstanding a sign convention here, but it’s not clear to me why H escape represents a loss of oxygen mass from the atmosphere system. The signs in equations (13) and (14) seem more intuitive. A word of explanation would be helpful.

The photochemical oxidation of CH<sub>4</sub> by O<sub>2</sub> in the atmosphere is represented as



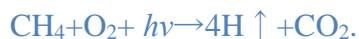
The sink of CH<sub>4</sub> and O<sub>2</sub> has a stoichiometric ratio of 1:2. However, when we consider hydrogen escape to space and following oxidation of C to CO<sub>2</sub>, the net reaction can be represented as (e.g., Goldblatt et al., 2006 *Nature*):



The sink of CH<sub>4</sub> and O<sub>2</sub> has a stoichiometric ratio of 1:1. Thus, the hydrogen escape to space halves the amount of O<sub>2</sub> consumed in the oxidation of CH<sub>4</sub>, representing a net O<sub>2</sub> production. In terms of the redox budget, the loss of hydrogen in equation (E2) represents the gain of the oxidizing power equivalent to one molecule of O<sub>2</sub>: 2H<sub>2</sub> + O<sub>2</sub> = 2H<sub>2</sub>O.

In equation (13), the net impact on the O<sub>2</sub> budget is represented, whereas internal source/sink fluxes of O<sub>2</sub> in the atmosphere are explicitly represented in equation (109). Both expressions are mutually consistent.

To clarify this point, we added a brief explanation of this to the revised manuscript (Line 1296-1299): “Note that the hydrogen escape to space is represented as the O<sub>2</sub> sink, because the hydrogen escape via CH<sub>4</sub> followed by the oxidation of carbon to CO<sub>2</sub> is represented as:



On the other hand, the photochemical oxidation of CH<sub>4</sub> can be written, as follows:



Thus, the hydrogen escape to space represents the net gain of oxidizing power to the system (see Eq. (13)).”

Finally, somewhere in the paper (perhaps in the introduction) it would be helpful to provide a brief history of the CANOPS model and its previous applications. This would help give the reader a better appreciation of the recent GRB improvements, as well as how the improved model might be applied.

We have added a brief explanation of this to the revised Introduction section (Line 110-114), and appreciate the opportunity to further clarify this point: “The core of this model is an ocean biogeochemical model, CANOPS (Ozaki et al., 2011; Ozaki and Tajika, 2013; Ozaki et al., 2019a). This model has been used to examine conditions for the development of widespread oceanic anoxia/euxinia during the Phanerozoic (Ozaki et al., 2011; Ozaki and Tajika, 2013; Kashiya et al., 2011) and to quantitatively constrain biogeochemical cycles during the Precambrian (Cole et al., 2022; Ozaki et al., 2019a; Ozaki et al., 2019b; Reinhard et al., 2017b).”

**Additional changes:**

- Various minor changes in wording and corrections throughout the manuscript.

We hope that we have adequately addressed all the suggestions raised by Reviewer 1, and appreciate their constructive feedback.