~~Robustness~~ The sensitivity of ~~simulating~~simulated aerosol climatic ~~impacts~~ impact to domain size using regional model~~: the sensitivity to domain size~~ (WRF-Chem v3.6)

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Key points:

1. Domain size has a great influence on the simulated meteorological fields and aerosol distribution during East Asian summer monsoon (EASM).
2. Regional simulations with different domain sizes demonstrate consistently that aerosols weaken EASM moisture transport.
3. Different domain sizes result in different strength of aerosol-induced changes of temperature and thus circulation and rainfall over China.

Abstract

Domain size can have significant ~~impacts~~impact on regional modeling results, but few studies ~~examining~~examined the sensitivities of ~~regional modeling results of~~simulated aerosol ~~impacts~~impact to regional domain size. This study investigates the regional modeling sensitivities of aerosol ~~impacts~~impact on East Asian summer monsoon (EASM) to domain size. The simulations with two different domain sizes demonstrate consistently that aerosols induce the cooling of lower~~-~~-troposphere that leads to the anti-cyclone circulation anomalies and thus the weakening of EASM moisture transport. The aerosol-induced adjustment of monsoonal circulation results in ~~a spatial~~an alternate increase and decrease pattern of ~~"+ + +" for~~ precipitation~~ change~~ over the continent of China. Domain size has a great influence on the simulated meteorological fields. For example, the simulation with ~~increasing~~larger domain size produces weaker EASM circulation, which also ~~affect~~affects aerosol distributions significantly. This leads to the difference of simulated strength and area extent of aerosol-induced changes of lower-tropospheric temperature and pressure, which further results in different ~~locations~~distributions of circulation and precipitation anomalies over the continent of China. For example, over Southeast China, aerosols induce the increase (decrease) of precipitation from the smaller-domain (larger-domain) simulation. Different domain sizes simulate consistently aerosol-induced increase of precipitation around 30°N over East China. This study highlights the important ~~impacts~~influence of domain size on regional modeling results of aerosol ~~impacts~~impact on circulation and precipitation, which may not be limited to East Asia. More generally, this study also implies that proper modeling of meteorological fields with appropriate domain size is one of the keys to simulate robust aerosol climatic ~~impacts~~impact.

1. Introduction

As one of the forcing's of climate change, aerosol contributes the largest uncertainty to the total radiative forcing estimate, and it has attracted more and more attention since the 1980s (IPCC, 2013; Li et al., 2019). ~~Aerosols~~Aerosol can absorb and scatter solar radiation through ~~Aerosol-Radiation~~aerosol-radiation interactions, affect the regional radiation budget, and amplify its impact through atmospheric mixing and circulation (e.g., Schwartz, 1996; Rinke et al., 2004; Kim et al., 2007; Z. ~~Q.~~ Li et al., 2010;~~ C.~~ Zhao et al., 2011, 2012, 2014; Myhre et al., 2013; Kuniyal et al., 2019; Zhang et al., 2020). Serving as cloud condensation nuclei or ice nuclei, ~~aerosols~~aerosol can change the microscopic and macroscopic characteristics of clouds and affect the climate, which is called ~~Aerosol-Cloud~~aerosol-cloud interactions (Twomey, 1977; Albrecht, 1989; Ackerman et al., 2000; Fan et al., 2012, 2013, 2016).~~ And there~~There are ~~many~~also some other possible ~~Aerosol-Cloud-Precipitation processes which~~aerosol-cloud-precipitation interactions that may amplify or dampen this effect (Rosenfeld et al., 2008, 2014; Tao et al., 2012; Fan et al., 2015, 2018).

Due to the large population and the rapid economic development in last few decades, East Asia has encountered large aerosol loading, and suffered from severe air pollution caused by various emission sources (e.g., Chan et al., 2008; X. Y. Zhang et al., 2012; Li et al., 2017; An et al., 2019). Moreover, East Asia is ~~located in~~within the monsoon region~~, the~~ and its weather and climate systems are more complicated, which makes the ~~study~~studying of aerosol effects more challenging (Ding et al., 2005; Ding, 2007; Li et al., 2016, 2019; Wu et al., 2016). In recent decades, the East Asian summer monsoon (EASM) and ~~summer~~the associated precipitation in eastern China have shown strong ~~interdecadal~~inter-decadal changes (Ding et al., 2008, 2013; Zhou et al., 2009; Zhu et al., 2011; Zhang, 2015), which ~~had~~has a significant impact on agriculture, economy, and human life (An et al., 2015). Many factors ~~are~~may be related to the ~~interdecadal~~inter-decadal variability of the EASM, such as extraterrestrial natural forcing, internal dynamical feedbacks within the climate system, and changes in atmospheric composition (e.g., greenhouse gases and aerosols) and surface conditions (land cover ~~changes~~change or urbanization) related to anthropogenic factors (Ding et al., 2008, 2009; H. ~~M.~~ Li et al., 2010; Song ~~&~~and Zhou, 2014; Xiao ~~&~~and Duan, 2016; Jiang et al., 2017). As one of the forcing factors~~ of summer climate change in East Asia~~, aerosol ~~have~~has attracted many people to study ~~the~~its effect on weather and climate ~~effects of summer aerosols in~~ of East Asia (Cowan ~~&~~and Cai, 2011; H. Zhang et al., 2012; Guo et al., 2013; Jiang et al., 2013, 2017; Wu et al., 2013; Song et al., 2014; Li et al., 2015, 2018; Wang et al., 2015, 2017; Chen et al., 2016; Kim et al., 2016; Xie et al., 2016; Dong et al., 2019).

Numerous studies have used global climate models to study the ~~impacts~~impact of anthropogenic ~~aerosols~~aerosol on the EASM climate and understand the mechanisms underneath (e.g., Guo et al., 2013; Jiang et al., 2013, 2017; Song et al., 2014; Yan et al., 2015; Chen et al., 2016; Wang et al., 2017; Li et al., 2018; Dong et al., 2019). The global modeling results have shown that ~~aerosols tend~~aerosol tends to reduce the land-sea thermal contrast, weaken the EASM, and thereby reduce the rainfall over the continent (e.g., Guo et al., 2013; Jiang et al., 2013; Song et al., 2014; Wang et al., 2017; Li et al., 2018; Dong et al., 2019). The reduction of monsoon precipitation over the continent may reduce the release of latent heat from condensation in the upper troposphere and further weaken the ~~East Asian summer monsoon~~EASM (e.g., Jiang et al., 2013; Li et al., 2019). Jiang et al. (2013) used the CAM5 (the Community Atmospheric Model version 5) model to study the effect of different aerosol types on ~~East Asian summer~~ clouds and precipitation during the EASM, and found that all anthropogenic aerosols suppressed the precipitation in North China and enhanced the precipitation in South China and the adjacent ocean areas. Through analyzing the CMIP5 (Coupled Model Intercomparison Program phase 5) modeling results, Song et al. (2014) examined the contributions of different ~~forcings~~forcing's (aerosol forcing, greenhouse gas forcing, and natural forcing) to the weakening of EASM circulation during 1958–2001~~, and~~. They found that aerosol forcing ~~plays~~played a major role ~~in the weakening of EASM, and the contribution of natural forcing is almost negligible, and the forcing of greenhouse gases is conducive to slightly strengthening rather than weakening the monsoon circulation~~.

Global climate models have been widely used for investigating aerosol ~~impacts, however~~impact. However, there are still large uncertainties with the results at regional scale, partly because the regional-scale monsoon rainband and aerosol distributions are still not able to be described accurately with relatively lower model horizontal resolution (H. ~~M.~~ Li et al., 2010; Guo et al., 2013; Jiang et al., 2013; Song et al., 2014; Li et al., 2018; Dong et al., 2019). In comparison, regional model often has relatively higher horizontal resolution and can better capture regional features of weather and climate systems and aerosol distributions, and therefore has been used to investigate aerosol regional climatic ~~impacts~~impact recently (e.g., Zhang et al., 2009; Stanelle et al., 2010; Zhao et al., 2011, 2012; Wu et al., 2013; Wang et al., 2015; Crippa et all, 2017; Zhuang et al., 2018). For example, ~~using~~Crippa et al. (2017) found that the enhanced resolution (from 60 to 12 km) can improve the regional model performance of meteorological fields and aerosol optical depth (AOD). Using the regional model

(RegCCMS), Wang et al. (2015) found that aerosol-cloud interaction ~~decreases~~decreased the ~~autoconversion~~auto-conversion rates of cloud water to rain water and ~~increases~~increased the liquid water path of clouds in East China, ~~strengthens~~which further strengthened the cooling of lower atmosphere caused by ~~the direct~~ aerosol-radiation ~~effect,~~interaction and ~~suppresses~~suppressed the convective precipitation. Wu et al. (2013), with the regional model (WRF-Chem), found that the aerosol heating effect caused the cloud to move northward over East China and led to the increased precipitation in the north.

Although regional model at higher horizontal resolution may better capture regional features of wind, cloud, precipitation, and aerosol, it also introduces additional uncertainties on modeling regional aerosol climatic ~~impacts~~impact resulted from the lateral boundary conditions of regional simulation. Previous studies have found that domain size of regional model ~~can~~could significantly influence the simulation results (e.g., Warner et al., 1997; Leduc and Laprise, 2009; Leduc et al., 2011; Bhaskaran et al., 2012; Diaconescu et al., 2013; Di Luca et al., 2015; Giorgi, 2019). For example, Bhaskaran et al. (2012) studied the sensitivity of ~~the~~ simulated hydrological cycle to the regional domain size over the Indian subcontinent. They found that the simulations with smaller domains produced the increased precipitation and evapotranspiration on seasonal mean and the higher number of moderate precipitation days relative to the ones with larger domains. Different distributions of cloud, precipitation, and winds from the simulations with different domain sizes may lead to different aerosol distributions and ~~its~~the associated climatic ~~impacts~~impact. Previous studies have found that aerosol ~~impacts~~impact on precipitation, clouds, and circulation ~~will~~would be significantly different under different weather and climate conditions (e.g., Wu et al., 2013; Wang et al., 2015; Xie et al., 2016). In addition, Seth and Giorgi. (1998) found that the smaller-domain simulation produced better precipitation compared with the observations, but resulted in an unrealistic response to the internal forcing. This indicates that the simulation domain size may also affect the aerosol ~~impacts~~impact on large-scale circulation. Therefore, the regional simulation with increased domain size may be preferred to ~~better~~ reflect the overall aerosol ~~impacts~~impact on large-scale circulation and ~~weather~~climate system without ~~the~~ strict constraint from ~~the~~ boundary forcing (e.g., Seth and Giorgi, 1998; Leduc and Laprise, 2009; Xue et al., 2014), but the increased domain size may make the ~~simulations~~simulation deviated from the forcing such as the reanalysis.

As far as we know, there are few studies examining the sensitivities of regional modeling results of aerosol ~~impacts~~impact to ~~regional~~ domain size. Although it can be expected that domain size will play a role, it is ~~not known~~unknown to what extent and how domain size can affect modeling results of aerosol climatic ~~impacts~~impact. Therefore, in this study, the regional online-coupled meteorology and chemistry model WRF-Chem (Weather Research and Forecasting model coupled with Chemistry) (Grell et al., 2005; Skamarock et al., 2008) is used to study the aerosol ~~impacts~~impact on ~~East Asian summer monsoon~~the EASM system ~~and~~with the focus on the modeling sensitivities to regional domain size. WRF-Chem has been widely used for studying aerosol meteorological and climatic ~~impacts~~impact over East Asia (e.g.,A. J. Ding et al., 2013; Wu et al., 2013; Gao et al., 2014; Chen et al., 2014; Zhao et al., 2014; Huang et al., 2016; Liu et al., 2016; Petaja et al., 2016; Zhao B et al., 2017). The investigation of aerosol ~~impacts~~impact under different simulated meteorological fields due to different domain sizes may also help understand the different modeling results about the aerosol ~~impacts~~impact on ~~East Asian summer monsoon~~EASM from previous studies. The study is organized as follows. Section 2 describes the numerical experiments and methods. The results and discussions are presented in Section 3. A summary is provided in Section 4.

2. Methodology

2.1 WRF-Chem

In this study, the version of WRF-Chem updated by the University of Science and Technology of China (USTC version of WRF-Chem) is used. The model simulates the emission, transport, mixing, and chemical transformation of trace gases and aerosols simultaneously with the meteorology, and can be used for investigation of regional-scale air quality and interactions between meteorology and chemistry. Compared with the publicly released version, the USTC version of WRF-Chem includes a few additional functions, such as the diagnosis of radiative forcing of aerosol species, optimized Kain-Fritsch (KF) convection scheme, aerosol-snow interaction, land surface coupled biogenic Volatile Organic Compound (VOC) emission, etc. (Zhao et al., 2013a, b, 2014, 2016; Hu et al., 2019; Du et al., 2020), all of which may have important impact on modeling aerosol and its climatic ~~impacts~~impact.

The Model for Simulating Aerosol Interactions and Chemistry (MOSAIC) aerosol module coupled with CBM-Z (carbon bond mechanism) photochemical mechanism in WRF-Chem is selected in this study (Zaveri & Peters, 1999; Zaveri et al., 2008). MOSAIC uses a sectional approach to represent aerosol size distributions with four or eight discrete size bins in the current version of WRF-Chem (Fast et al., 2006). To reduce the computational cost, four discrete size bins is selected in this study. All major aerosol components including sulfate, nitrate, ammonium, black carbon, organic matter, sea-salt, mineral dust, and other inorganic matter (OIN) are simulated in the model. The MOSAIC aerosol scheme includes physical and chemical processes of nucleation, condensation, coagulation, aqueous-phase chemistry, and water uptake by aerosols. Dry deposition of aerosol mass and number is simulated following the approach of Binkowski and Shankar (1995), which includes both turbulent diffusion and gravitational settling. Wet removal of aerosols by grid-resolved stratiform clouds and precipitation includes in-cloud removal (rainout) and below-cloud removal (washout) by impaction and interception, following Easter et al. (2004) and Chapman et al. (2009). In this study, cloud-ice-borne aerosols are not explicitly treated in the model, but the removal of aerosols by the droplet freezing process is considered. Convective transport and wet removal of aerosols by cumulus clouds is coupled with the Kain-Fritsch cumulus scheme as Zhao et al. (2013b). Aerosol radiative feedback is coupled with the Rapid Radiative Transfer Model (RRTMG) (Mlawer et al., 1997; Iacono et al., 2000) for both SW and LW radiation as implemented by Zhao et al. (2011). The optical properties and direct radiative forcing of individual aerosol species in the atmosphere are diagnosed following the methodology described in Zhao et al. (2013a).

2.2 Numerical experiments

Four sets of experiments, CTRL-L, CTRL-S, CLEAN-L, and CLEAN-S, with different simulation domain sizes or emission configurations are conducted as explained and listed in Table 1. The control experiments (CTRL-S~~,~~ and CTRL-L)~~ simulations~~ use standard anthropogenic emission dataset (described in Section 2.3), while the clean simulations (CLEAN-S~~,~~ and CLEAN-L) apply a factor of 0.1 on the standard emissions within the small domain to represent a clean atmosphere condition over East Asia (Fig. 1). The CTRL-L and CTRL-S (CLEAN-L and CLEAN-S) represent the simulations with large and small domain sizes, respectively, as shown in ~~Figure~~Fig. 1. The aerosol ~~impacts~~impact can be calculated by the difference between the CTRL and CLEAN simulations for each simulation domain. The comparison of aerosol ~~impacts between the large and small simulation domains implies the sensitivity of aerosol impacts to domain size.~~ impact between the large and small simulation domains implies the sensitivity of aerosol impact to domain size. Besides these experiments, another set of experiment NoRA-S is conducted to isolate aerosol-radiation and aerosol-cloud interactions for further understanding the mechanisms of aerosol impact, which is also listed in Table 1. The horizontal resolution of 30 km is selected for both simulation domains with the consideration of the balance of computational efficiency and modeling performance, particularly for the larger domain. The comparable horizontal resolutions have also been widely used for investigating aerosol impact on regional climate (e.g., Zhang et al., 2009; Stanelle et al., 2010; Zhao et al., 2011, 2012; Chen et al., 2014; Wang et al., 2015).

All the WRF-Chem experiments select the Morrison two-moment microphysics (Morrison et al., 2009), Kain-Fritsch cumulus scheme (Kain, 2004), unified Noah land-surface model, Rapid Radiative Transfer Model (RRTMG) longwave and shortwave radiation schemes (Iacono et al., 2008), and MYNN planetary boundary layer (PBL) scheme (Nakanishi & Niino, 2006,2009). Following Du et al. (2020), the PBL mixing coefficient is modified to simulate better PBL mixing of aerosols. Five ensemble simulations are performed for each experiment by changing the simulation initial ~~conditions~~time at UTC 0000 from May ~~12~~12th to May ~~16~~16th, 2017~~.~~ (i.e., the five ensemble simulations start at UTC 0000 of May 12th, 13th, 14th, 15th, and 16th, respectively). The averaged results from five ensembles are analyzed to reduce the influence of modeling internal variability. The simulations run continuously through entire June and July of 2017. The analysis focuses on the simulation results for June 1 to July 31, 2017. The meteorological initial and lateral boundary conditions are derived from National Centers for Environmental Prediction (NCEP) Final (FNL) Operational Global Analysis data (NCEP, 2000) with a horizontal resolution of 1°×1° and a ~~time~~temporal resolution of 6h. The chemical initial and boundary conditions are provided by a quasi-global WRF-Chem simulation for the same time period. The quasi-global WRF-Chem simulation is performed at 1°×1° horizontal resolution with 360×130 grid cells (180°W-180°E, 60°S-70°N). More details about the general configuration of a quasi-global WRF-Chem simulation can be found in Zhao et al. (2013b) and Hu et al. (2016). The simulation configuration is summarized in Table 2.

2.3 Emissions

Biomass burning emissions are obtained from the Fire Inventory (FINN) of the National Center for Atmospheric Research with hourly temporal resolution and 1 km horizontal resolution (Wiedinmyer et al., 2011), and the injection heights follow Dentener et al. (2006) for the Aerosol Comparison between Observations and Models (AeroCom) project. The natural dust emission fluxes are calculated based on the adjusted GOCART dust emission scheme (Ginoux et al., 2001; Zhao et al., 2010), and the emitted dust particles are distributed into the MOSAIC aerosol size bins following a theoretical expression based on the physics of scale-invariant fragmentation of brittle materials derived by Kok (2011). More details about the dust emission scheme coupled with MOSAIC aerosol scheme in WRF-Chem can be found in Zhao et al. (2010, 2013b). Sea-salt emission follows Zhao et al. (2013a), which includes the correction of particles with radius less than 0.2 μm and the dependence of sea-salt emission on sea surface temperature. Anthropogenic emissions are obtained from the Multi-resolution Emission Inventory for China (MEIC) at 0.1~~°x0~~°×0.1° horizontal resolution and with monthly temporal resolution for 2015 (Li et al., 2017; Zheng et al., 2018), except that the emissions outside of China are from the Hemispheric Transport of Air Pollution version2 (HTAPv2) at 0.1~~°x0~~°×0.1° horizontal resolution and with monthly temporal resolution for the year 2010 (Janssens-Maenhout et al., 2015) (Fig. 1). As discussed above, the anthropogenic ~~emissions~~emission in the CLEAN experiments ~~is~~has a factor of 0.1 of that in the CTRL experiment~~, and~~. In addition, in the CLEAN-L experiment, only the emissions in the area of ~~the~~ small domain (denoted by the red box) are adjusted. In this way, the emission reduction from the simulations with both domains are made consistent.

2.4 Observations and reanalysis

Although the aims of this study are not evaluating the simulation results to determine the optimal model configuration for the experiments, some observations and reanalysis datasets are still used to provide the references for the key fields. The comparison with these references can demonstrate whether the simulation results are acceptable for further analysis. The MISR (Multi-angle Imaging SpectroRadiometer, instrument on board the NASA Terra platform) retrieval dataset is used as a reference of spatial distribution of AOD (Diner et al, 1998; Martonchik et al., 2004). When showing the comparison between the MISR retrieved and the simulated AOD, the simulation results are sampled from 10 am - 11 am for averaging and at the locations of the retrievals because the Terra platform passes over the equator at about 10:45 LT (Diner et al, 2001). The precipitation datasets of CMA (National Meteorological Information Center of China Meteorological Administration) and CMORPH (Climate Prediction Center MORPHing technique) are used as the references for spatial and temporal variations of precipitation during the simulation period. The CMORPH dataset is a global precipitation reanalysis dataset that is derived from geostationary satellite IR imagery (Joyce et al., 2004). The CMA rainfall was measured by tipping buckets, self-recording siphon rain gauges, or automatic rain gauges and was subject to strict quality control. The European Centre for Medium-Range Weather Forecasts (ECMWF) Reanalysis v5 (ERA5) are used as a reference for wind fields (Hersbach, 2020).

3. Results

3.1 Sensitivity of simulated meteorological fields to domain size

Figure 2 shows the spatial distributions of precipitation and moisture transport at 700 hPa over the small domain averaged for June and July of 2017 from the observation and reanalysis, and the simulations of CLEAN-S and CLEAN-L. The observation and reanalysis show that the southwesterly transports large amount of moisture into East China. The converge of large amount of moisture results in heavy precipitation over southern China and its adjacent ocean. Due to the gradual weakening of northeastward moisture transport and the blocking effect of ~~the~~ western mountains, precipitation becomes much weaker over northern and western China. Compared with the CMORPH observation and ERA5 reanalysis (Fig. 2), CLEAN-S can reasonably produce the spatial ~~distribution~~distributions of precipitation and moisture transport at 700 hPa, with slight underestimation of meridional moisture transport over eastern China. It is evident that the meridional moisture transport over southern China becomes weaker with the increasing domain size, and the eastward transport becomes stronger. In addition, the overall southwesterly moisture transport ~~shift~~shifts to the east. This leads to a decrease of precipitation over eastern China and an increase over the East China Sea. Compared with the observations of hourly precipitation from the CMA stations over eastern China (Fig. S1 in the supporting material), both the CLEAN-S and CLEAN-L experiments can generally reproduce the daily variation of precipitation over eastern China, although the CLEAN-L simulated precipitation is lower consistent with its weaker moisture transport over the region.

The difference in moisture transport between the simulations with different domain sizes results from their difference in geopotential height and wind circulation. Figure 3 shows the spatial distributions of geopotential height (GPH) and wind field at 700 hPa from the ERA5 reanalysis and the CLEAN-S simulation, and of the difference between CLEAN-L and CLEAN-S. The comparison with the ERA5 reanalysis shows that the CLEAN-S can well simulate the distributions of GPH and wind fields at 700 hPa ~~(Fig. S2 in the supporting material).~~. The spatial distribution of wind fields is generally consistent with that of moisture transport (Fig. 2~~),~~) and is largely controlled by the West Pacific sub-tropical high (WPSH). Compared to CLEAN-S, CLEAN-L simulates lower GPH at 700 hPa and produces an anomalous lower pressure center on the East China Sea, which indicates the weaker WPSH with increasing domain size. This causes the southwestward wind anomalies over the continent, which weakens the monsoon driven northeastward moisture transport. Over the South China Sea, the westerly anomalies enhance the eastward transport of moisture.

The impact of domain size is not only on the horizontal distribution of wind fields but also on the vertical circulation. Figure 4 shows the cross-section of meridional temperature anomalies and wind averaged for 105°E and 122°E from the ERA5 reanalysis and the CLEAN-S simulation during June to July, and of the difference of temperature (not meridional temperature anomalies) and wind between CLEAN-L and CLEAN-S. The meridional temperature anomalies are calculated by subtracting the mean temperature in this latitude range (as shown in Fig. 4) at each pressure level. First of all, CLEAN-S can ~~general~~generally reproduce the temperature gradient and wind circulation from the ERA5 reanalysis~~ (Fig. S3 in the supporting material).~~. Relatively large meridional temperature gradient exists between 700 hPa and 200 hPa, where the temperature is higher over the South. Below 700 hPa, the temperature gradient is relatively weaker~~,~~ and the temperature is higher over the North. Along with this distribution of temperature gradient, meridional wind blows from the South and the North and converges at the latitude around 34°N, which generates strong upward motion in the area of 20°N-35°N. This is consistent with the spatial ~~distribution~~distributions of precipitation and moisture transport (Fig. 2). Compared with the CLEAN-S experiment, the CLEAN-L experiment produces larger meridional temperature gradient between 700 hPa and 200 hPa and weaker gradient below 850 hPa. The circulation from~~ the~~ CLEAN-L is generally consistent with CLEAN-S, but the southerly wind from CLEAN-L is weaker and the northerly wind is stronger. This results in an overall northerly wind anomalies from CLEAN-L compared with CLEAN-S, ~~and also~~along with a southward shift of~~ the~~ wind convergence from 34°N to 32°N. It is also noteworthy that the upward motion is weakened around 22°N-38°N and strengthened to the south of 20°N due to the increased domain size.

3.2 Sensitivity of simulated aerosol characteristics to domain size

Figure 5 shows the spatial ~~distribution~~distributions of AOD averaged for June and July of 2017 from the CTRL-S simulation, and of the difference between CTRL-L and CTRL-S. ~~It can be seen that relatively~~Relatively high AOD (>0.6) exists in the Sichuan Basin and the North China ~~Plain. The~~plain. AOD over East Central China and South China is relatively lower (0.2-0.5), which is in line with previous ~~research (~~studies (e.g., Luo et al., 2014; Qi et al., 2013). In general, ~~the ~~CTRL-S ~~generally ~~captures the spatial distribution of retrieved AOD from MISR (Fig. ~~S4~~S2 in the supporting material). Compared with the CTRL-S experiment, CTRL-L simulates a similar spatial pattern of AOD ~~as CTRL-S, ~~but produces higher AOD in southern China and lower AOD in most areas of northern China. To explore the reasons of difference between the two simulations, Figure 6 shows the spatial distributions of column integrated ~~total~~ $PM_{2.5}$ ~~concentration~~mass and water content in aerosol averaged ~~for~~in June and July of 2017 ~~from the CTRL-S simulation, and~~and of the difference between CTRL-L and CTRL-S. The

CTRL-S simulation shows high $PM_{2.5}$ mass loading over the North China ~~Plain~~plain, which is consistent with the spatial distribution of AOD (Fig. 5). The $PM_{2.5}$ mass loading also shows high values over Northwest China, which is not shown in the spatial distribution of AOD. This is mainly due to the high mass loading of dust over Northwest China (Fig. ~~S5~~S3 in the supporting material~~),~~) and the water content associated with dust is relatively small.

CTRL-L simulates higher $PM_{2.5}$ mass loading over Southeast China and lower values over North China, which is consistent with AOD. The difference of water content in aerosol shows a similar pattern. The analysis shows that the difference of $PM_{2.5}$ mass loading over North China is mainly due to the difference of dust, while the difference over Southeast China is due to anthropogenic aerosols (Fig. ~~S5~~S3). The reduction of dust mass loading over North China from CTRL-L is primarily due to its weakening of westerlies over Northwest China compared to CTRL-S (Fig. 3), which results in less transport of dust into the downwind region. The increase of aerosol mass loading over Southeast China in CTRL-L is partly due to its less wet scavenging associated with weak precipitation (Fig. 2). The weakening of northward transport of aerosol (Fig. 3) also contributes to the increase of $PM_{2.5}$ mass loading over southern China in CTRL-L. Besides the change of dry aerosol mass loading, the change of water content in aerosol between the two experiments also contributes to the change in AOD, which results from the difference of both dry aerosol mass and moisture.

Figure 7 shows the latitude-height cross-section of total $PM_{2.5}$ averaged between 105°E and 122°E for June and July of 2017 from the CTRL-S experiment, and of the difference between CTRL-L and CTRL-S. The latitudinal distribution of aerosol is consistent with ~~it~~its spatial pattern with high aerosol mass concentration over North China. The mass concentration gradually reduces from the surface to the free atmosphere. The mass concentration around 500 hPa over North China can reach 5 ug/m^3 that is comparable to the surface concentration over South China. In general, CTRL-L simulates higher aerosol mass concentration over South China and lower aerosol mass concentration over North China from the surface to about 500 hPa. At 32°N-36°N, CTRL-L simulates lower aerosol mass concentration near the surface and higher ~~above~~ values between 700 hPa and 850 hPa~~.~~, likely due to the difference in aerosol wet scavenging and transport between the two experiments. The difference ~~of aerosol~~in horizontal and vertical distributions of aerosols and also the circulation patterns between the two experiments may lead to the difference in simulating aerosol ~~impacts on East Asian monsoon system~~impact on EASM.

3.3 Sensitivity of aerosol impact to domain size

Before studying the sensitivity of aerosol ~~impacts~~impact to domain size, the ~~impacts~~impact of aerosol on precipitation and circulation from the small domain ~~simulations are~~experiments is first investigated. Figure 8 shows the spatial distributions of aerosol-induced difference (CTRL-CLEAN) of precipitation and moisture transport at 700 hPa averaged for June and July of 2017 from the small domain simulations. The dominant effect is that aerosol weakens the southwesterlies ~~flow~~ and reduces the moisture transport over the continent of Central and South China (primarily between 105°E-115°E). Along the coast of Southeast China, the moisture transport is enhanced slightly. Over the continent of China, aerosol induces ~~a "+-+-+" pattern~~an alternate increase and decrease pattern (denoted as "+-+-+") of precipitation changes, i.e., precipitation increases in the south of 25°N, north of 40°N, and around 30°N, while decreases at 25°N~30°N and 32°N~40°N. This weakening of monsoonal circulation at the lower troposphere is found mainly due to the cooling of lower troposphere and thus the increase of surface pressure by aerosols (Fig. 9). The temperature averaged for lower-troposphere (below 500 hPa) is reduced by aerosols over the continent of China, which results in a positive pressure anomaly center in Southwest China. This leads to an anticyclone anomaly as shown in Fig. 8, which weakens the monsoonal southwesterlies between 105°E-115°E.

In order to further understand the mechanisms of aerosol ~~impacts~~impact and isolate aerosol-radiation and aerosol-cloud interactions, another set of numerical experiment (NoRA-S) with the small domain ~~are~~is conducted, similar as CTRL-S but with the aerosol-radiation interaction turned off. The difference of results between NoRA-S and CLEAN-S (NoRA-S minus CLEAN-S) is interpreted as the ~~impacts~~impact of aerosol-cloud interaction, while the difference of results between CTRL-S and NoRA-S (CTRL-S minus NoRA-S) is interpreted as the ~~impacts~~impact of aerosol-radiation ~~interactions~~interaction. Figure 10 shows the spatial ~~distribution~~distributions of the ~~impacts~~impact of aerosol-cloud and aerosol-radiation interactions on (a, d) tropospheric temperature averaged below 500 hPa, (b, e) surface pressure, (c, f) precipitation and moisture transport. The aerosol-cloud interaction reduces significantly the lower-tropospheric temperature (Fig. 10a) over a large area of South China (to the south of 32°N) mainly due to its increasing of cloud amounts (Fig. ~~S6a~~S4a in the supporting material) over this area, which results in an increase of surface pressure in this area (Fig. 10b). Similarly, aerosol-cloud interaction also increases cloud amounts over Northeast China and its adjacent ocean (Fig. ~~S6a~~S4a) and thus reduces the lower-tropospheric temperature and increases the surface pressure over the area. The surface pressure over the Yellow River Basin is reduced slightly by aerosol-cloud interaction, which may be due to the reduction of cloud amounts (Fig. ~~S6a~~S4a) and the increase of lower-tropospheric temperature. Although, the experiments can generally demonstrate that aerosol-cloud interaction can largely affect cloud amount, lower-tropospheric temperature, and surface pressure, please note that the co-locations of the changes of cloud, temperature, and surface pressure may not be simply straightforward. For example, in a fully coupled system, the cloud change due to aerosols would also adjust the temperature through the release of latent heat in the atmosphere. In addition, the change of temperature would also modulate the circulation and further feedback to the distributions of cloud and temperature. The difference between NoRA-S and CLEAN-S over Northwest China is due to the dust-radiation interaction that is included in CLEAN-S but not in NoRA-S. The analysis of this study focuses on the ~~impacts~~impact of anthropogenic aerosol. The combined effect of two anti-cyclone anomalies due to the two positive pressure anomalies at the lower~~-~~-troposphere results in the southward wind anomalies over the ocean and the northward wind anomalies over North China, while the changes of circulations in other areas of China is negligible.

The primary ~~impacts~~impact of aerosol-radiation interaction on lower-atmospheric temperature ~~are~~is the positive temperature ~~anomalies~~anomaly over the Yellow Ocean and over central China and the negative temperature ~~anomalies~~anomaly over the Yellow River Basin and Southwest China, which is the combined ~~effects~~effect from the aerosol cooling and heating at the surface and ~~heating~~ in the atmosphere, respectively, and ~~also~~ the adjustment of cloud distributions (Fig. ~~S6b~~S4b and Fig. ~~S7~~S5). The two positive temperature anomaly centers lead to two negative pressure anomaly centers and thus a large cyclone circulation anomaly over the continent of East China. Therefore, it can be noted that the influence of aerosol-cloud and aerosol-radiation interactions on monsoonal circulations ~~are~~is counteracted over the ocean and over northern China, which results in relatively small changes of monsoonal circulation over the ocean and over northern China (Fig. 8). The overall aerosol impact is shown as the weakening of the monsoonal circulation over the continent of central and southern China (Fig. 8), which is mainly contributed by the aerosol-radiation interaction.

Figure 11 shows the latitude-pressure cross-section of aerosol-induced difference (CTRL-CLEAN) of temperature and wind averaged between 105°E and 122°E for June and July of 2017 from the small domain simulation. ~~It can be seen that the~~The pattern of precipitation change corresponds well to the ~~changes~~change of wind circulation. The weakening of monsoonal southwesterlies ~~result~~results in a sinking airflow anomaly around 28°N and the compensating upward anomaly around 24°N in the south of China, and also a downdraft around 35°N and an updraft around 40°N in north China. These two sinking airflows ~~corresponds~~correspond to the reduced precipitation between 25°N and 30°N and between 32°N and 40°N, respectively (Fig. 8), while these updrafts correspond to the increasing precipitation between 22°N and 25°N and between 32°N and 40°N. There is also weak upward compensating airflow around 30°N, leading to the slight increase of precipitation in the area (Fig. 8). It is noteworthy that aerosols lead to an abnormal cooling center around 33°N between 400 hPa to 200 hPa. This is mainly because of less solar radiation entering the atmosphere due to aerosol-radiation and aerosol-cloud interactions, and also weaker monsoonal airflow that leads to less release of latent heat from cloud and precipitation (Fig. ~~S8~~S6 in the supporting material). This cooling anomaly center also strengthens the downdraft anomalies on its both sides, further weakening the monsoonal circulation.

In order to explore the sensitivity of aerosol ~~impacts~~impact to domain size, similar as Fig. 8, Figure 12 shows the results from the large domain simulations. One consistent signal between the simulations with different domain sizes is that aerosols weaken the southwesterlies ~~flow~~ and reduce the moisture transport over the continent of Central and South China. The difference is that this weakening is not only over the inland of China but also extending to over the South China Ocean. The weakening of monsoon airflow is broader with the increasing domain size, which may be due to its weaker monsoon airflow (Fig. 3) and less constraint from the lateral ~~boundaries~~boundary conditions in the large domain simulation. Another consistent signal between the two sets of simulations with different domain sizes is that aerosol induces a similar "+-+-+" pattern of precipitation changes over the domain, except that the areas with precipitation reduction become broader. This leads to the precipitation reduction over almost the entire region between 20°N~40°N over the continent of China except the area around 30°N with increasing precipitation. The increases of precipitation on the two sides of precipitation reduction area shift southward to the South China ocean and northward to the north of 40°N, respectively.

Similar as the small domain simulation, the weakening of monsoonal airflow in the large domain simulation is also due to the abnormal positive lower-level pressure that is caused by the lower~~ atmosphere~~-tropospheric cooling (Fig. 13), which can also be explained by the effects of aerosol-radiation and aerosol-cloud interactions (Fig. ~~S9~~S7 and Fig. ~~S10~~S8 in the supporting material). However, compared with the small domain simulation (Fig. 9), the cooling anomaly of lower-tropospheric temperature and thus the positive anomaly of lower-level pressure ~~cover~~covers a broader area from the large domain simulation. The two aerosol-induced cooling centers over the continent of China lead to two positive lower-level pressure anomaly centers that ~~results~~result in a large anti-cycle circulation anomaly (Fig. 12), which weakens the monsoonal southwesterly airflow over South China and the South China Ocean and also slightly enhances the southwesterly over West China. Again, the pattern of precipitation change corresponds well to the ~~changes~~change of wind circulation (Fig. 14). With larger domain size, aerosols lead to a broader area (between 20°N~40°N) of abnormal cooling in the troposphere up to 200 hPa. The single cooling center in the small domain simulation is split into two centers, one around 30°N at 250 hPa and another around 36°N at 700 hPa. The weakening of ~~the~~ background circulation and broader cooling area lead to the broader sinking ~~airflows~~airflow over the region, which results in the broader area of reduced precipitation compared with the small domain simulation (Fig. 8 and Fig. 12). The increasing precipitation around 30°N is also resulted from the compensating ~~updrafts~~updraft around 30°N.

4. Summary and Discussion

Due to the importance of domain size on regional modeling results and few studies ~~examining~~that examined the sensitivities of regional modeling results of aerosol ~~impacts~~impact to domain size, this study applies the WRF-Chem model to simulate the anthropogenic aerosol ~~impacts~~impact on ~~East Asian summer monsoon~~EASM circulation and precipitation, focusing on the modeling sensitivities to regional domain size. The ~~impacts~~influence of domain size on meteorological fields, aerosol characteristics, and aerosol ~~impacts are~~impact is investigated.

First of all, the domain size has a great influence on the simulated meteorological fields. From the ~~small~~smaller domain simulation, the circulation and precipitation are in good agreement with the reanalysis data and observations. The ~~large~~larger domain simulation produces weaker ~~East Asian summer monsoon system~~and southward shifting ~~southward~~EASM system, which results in that the precipitation ~~decrease~~decreases in southern China and ~~increase~~increases in the adjacent ocean. The changes of circulation and precipitation also lead to the increase of aerosol mass loading in southern China and the decrease in northern China in the ~~large~~larger domain simulation. The deviation of atmospheric fields particularly the circulation between the simulations with different domains is partly due to their different constraint from lateral boundary conditions. With the less constraint of ~~the~~ boundary forcing ~~from,~~ the ~~reanalysis data, the large~~larger domain simulation may produce negative bias in precipitation over the Yangtze River Basin and positive bias in ~~water vapor~~moisture transport over the South China Ocean as reported by previous studies. The uncertainties in moisture transport prescribed in the lateral ~~boundaries~~boundary conditions from the reanalysis over a larger domain may also contribute to the biases (e.g., Wang and Yang, 2008; Huang and Gao, 2018). ~~With~~Previous studies found that, with the larger ~~domain, the~~ simulation ~~includes larger area~~domain including more areas of ocean~~. Without~~, without considering the ~~online~~ interaction between the atmosphere and the ocean (i.e., with prescribed SST from the reanalysis), the artificial positive feedback between precipitation and surface latent heat flux may overestimate the precipitation over the subtropical Western North Pacific (WNP) and inhibit the westward expansion of the WNP subtropical high (e.g., Cha and Lee, 2009; Lee and Cha, 2020).

In terms of the climatic ~~impacts~~impact of anthropogenic aerosols on ~~East Asian summer monsoon~~EASM, as shown in the schematic ~~plot~~figure (Fig. 15), aerosols induce the cooling of lower troposphere over the continent through aerosol-radiation and aerosol-cloud interactions, which leads to an increase of regional pressure at lower atmosphere. The regional positive pressure anomalies result in the anti-cyclone circulation anomalies and thus ~~weakens~~weaken the summer monsoonal northeastward moisture transport, which is consistent with previous studies (e.g., Y. ~~Q.~~ Jiang et al., 2013; Song et al., 2014; T.~~ J.~~ Wang et al., 2015; Xie et al., 2016). The weakening of monsoonal circulation leads to several sinking airflows and compensating updrafts that correspond well to the regions with the decrease and increase of precipitation, respectively, showing a spatial pattern of "+-+-+" for precipitation change. The difference in ~~the~~ aerosol ~~impacts~~impact from the numerical experiments with different domain sizes is mainly determined by their simulated different strength and area extent of the aerosol-induced lower-tropospheric negative temperature anomalies. Compared with the smaller-domain simulation, the larger-domain simulation with weaker monsoonal circulation generates a broader area with negative temperature and positive pressure anomalies at the lower troposphere, which results in broader sinking airflows and thus broader areas of precipitation reduction over the continent of China. This could lead to the opposite signals of precipitation change due to aerosols over China~~, for~~. For example, over Southeast China, the ~~increase and decrease of~~ precipitation is increased (decreased) from the smaller-domain ~~and~~ (larger-domain ~~simulations, respectively.~~) simulation. The consistent signal of aerosol ~~impacts~~impact between the simulations with different domain sizes is the increasing precipitation around 30°N that is resulted from the compensating updraft over the region.

Although the modeling results of aerosol ~~impacts~~impact in this study may have some uncertainties associated with physical and chemical processes, emissions, and ~~simulation~~ horizontal resolutions~~,~~ (e.g., Di Luca et al., 2015; Crippa et al., 2019), it highlights the ~~impacts~~impact of simulation domain size on regional modeling aerosol ~~impacts~~impact on monsoonal circulation and precipitation, which may not be limited to the region of East Asia. Uncertainties of modeling aerosol climatic ~~impacts~~impact are often investigated ~~focusing~~with the focus on aerosol characteristics such as their distributions and properties. This study adds another complexity (impact of domain size) on regional modeling of aerosol climatic

~~impacts~~impact. More specifically, ~~although~~on one hand, larger-domain simulation may better allow aerosol feedbacks~~ of aerosol impacts~~ on weather and climate systems without strong lateral boundary constraint (e.g., Seth and Giorgi, 1998; Leduc and Laprise, 2009~~),~~; Diaconescu et al., 2013), but it may produce biased meteorological fields compared to smaller-domain simulation, which can then significantly influence the modeling results of aerosol ~~impacts.~~impact. On the other hand, although the simulation with smaller domain produces better large-scale circulation compared to the reanalysis, the lateral boundary condition may also have stronger constraint on aerosol feedbacks to large-scale circulation. Therefore, not like meteorological fields or aerosol properties, there is no direct observation or reanalysis that can be used as the references to evaluate aerosol impact (Di Luca et al., 2015; Crippa et al., 2017), and the optimal configuration of simulation domain is hard to be determined in this study. It may be the key to ~~simulate reasonable/less biased~~improve the simulated meteorological fields with larger regional domain or global domain in order to model robust aerosol climatic ~~impacts~~impact. More generally, this study also highlights the ~~impacts~~impact of background meteorological fields (without aerosol effect) on simulated aerosol ~~impacts~~impact. Proper modeling of background meteorological fields is one of the keys to simulate ~~robust~~reliable aerosol climatic ~~impacts~~impact. The model inter-comparison study of aerosol climatic ~~impacts~~impact should also focus on the diversity of simulated background meteorological fields besides aerosol characteristics.

~~Data statement~~

Code and data availability

The release version of WRF-Chem can be downloaded from http://www2.mmm.ucar.edu/wrf/users/download/get_source.html. The code of updated USTC version of WRF-Chem is available at https://doi.org/10.5281/zenodo.4663508 or contact ~~chunzhao@ustc.edu.cn.~~chunzhao@ustc.edu.cn. The dataset from the European Centre for Medium-Range Weather Forecasts (ECMWF) Reanalysis v5 (ERA5) can be downloaded from https://rda.ucar.edu/datasets/ds633.1/ (last access: Aug 2021). The CMORPH data can be downloaded from https://ftp.cpc.ncep.noaa.gov/precip/CMORPH_V1.0/CRT/0.25deg-DLY_00Z/2017/ (last access: Aug 2021).

Author contributions

Xiaodong Wang and Chun Zhao designed the experiments, conducted and analyzed the simulations. All authors contributed to the discussion and final version of the paper.

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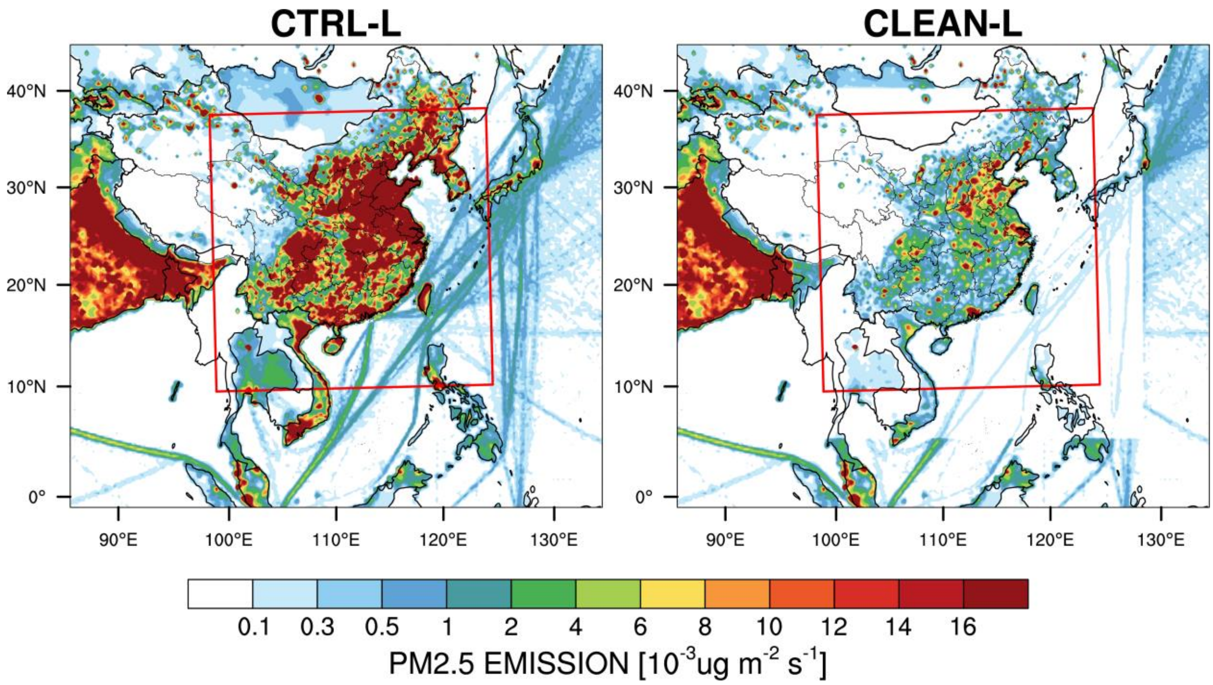
Table 1. Experiment Description.

Experiment ID	Experiment Description
CTRL-L	Control experiment with large simulation domain.
CLEAN-L	~~Same~~Similar as CTRL-L, but the anthropogenic aerosol emissions are 0.1 times of CTRL-L.
CTRL-S	Control experiment with small simulation domain.
CLEAN-S	~~Same~~Similar as CTRL-S, but the anthropogenic aerosol emissions are 0.1 times of CTRL-S.
NoRA-S	Similar as CTRL-S, but with the aerosol-radiation interaction turned off.

Table 2. Summary of model configurations.

Description	Selection (L, S)
Horizontal grid spacing	30km
Grid dimensions	201x231, 121x121
Vertical layers	41
Topography	USGS_30s
Model top press	100hPa
Aerosol scheme	MOSAIC 4 bin
Gas-phase chemistry	CBM-Z
Long wave Radiation	RRTMG
Short wave Radiation	RRTMG
Cloud Microphysics	Morrison 2-moment
Cumulus Cloud	Kain-Fritsch
Planetary boundary layer	MYNN 3rd
Land surface	unified Noah land-surface model
Meteorological Forcing	FNL, 1°x1° ,6 hourly


CTRL-L
CLEAN-L
40°N
30°N
20°N
10°N
0°
90°E
100°E
110°E
120°E
130°E
0.1 0.3 0.5 1 2 4 6 8 10 12 14 16
PM2.5 EMISSION [10^{-3}ug m^{-2} s^{-1}]

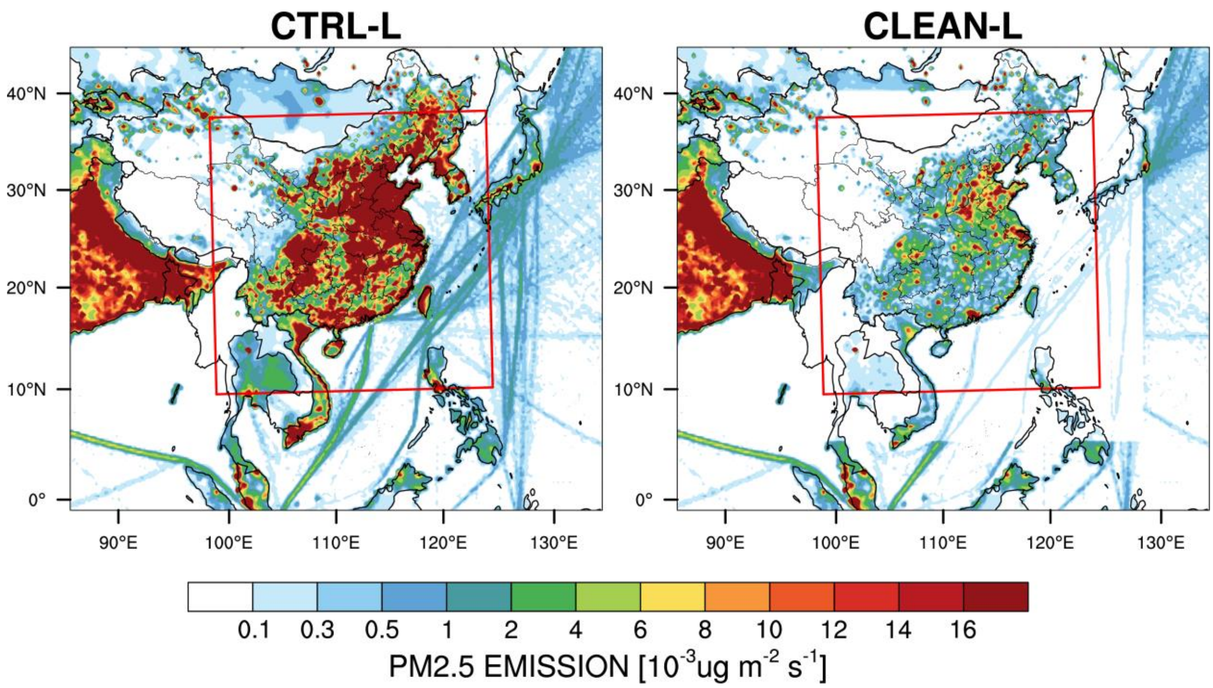
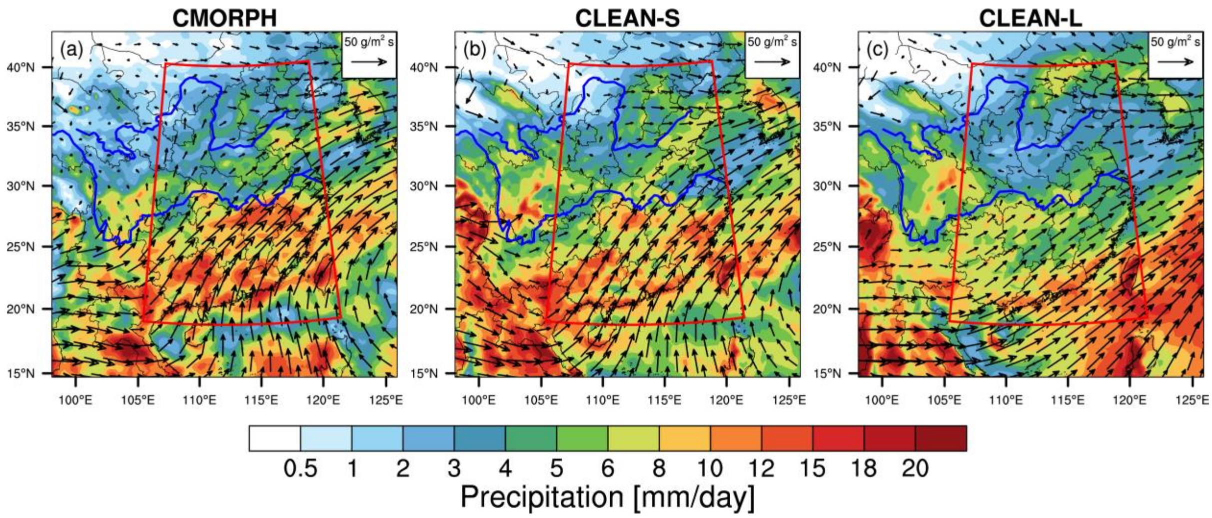

CTRL-L
CLEAN-L
40°N
30°N
20°N
10°N
0°
90°E
100°E
110°E
120°E
130°E
0.1 0.3 0.5 1 2 4 6 8 10 12 14 16
PM2.5 EMISSION [10^{-3}ug m^{-2} s^{-1}]

Figure 1. Spatial distributions of anthropogenic emissions of primary $PM_{2.5}$ averaged for June and July for the simulation domains. The red box in the large simulation domain represents the small domain.



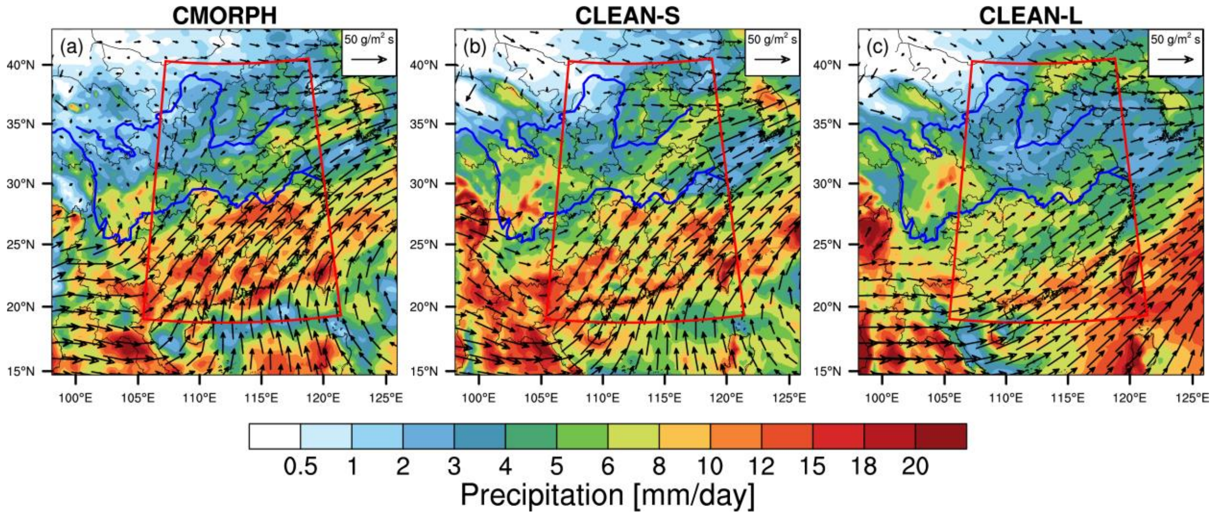


Figure 2. Mean precipitation rate (mm/day) and 700hPa moisture transport (g/m² s) over the small domain for the two months of June and July 2017 from (a) CMORPH and ERA5 reanalysis, (b) CLEAN-S simulation, and (c) CLEAN-L simulation. The red box (20°N-42°N, 105°E-122°E) represents the focus area of analysis in follow. (a) Precipitation data comes from CMORPH, and the 700hPa moisture transport field data is obtained by processing ERA5 reanalysis.

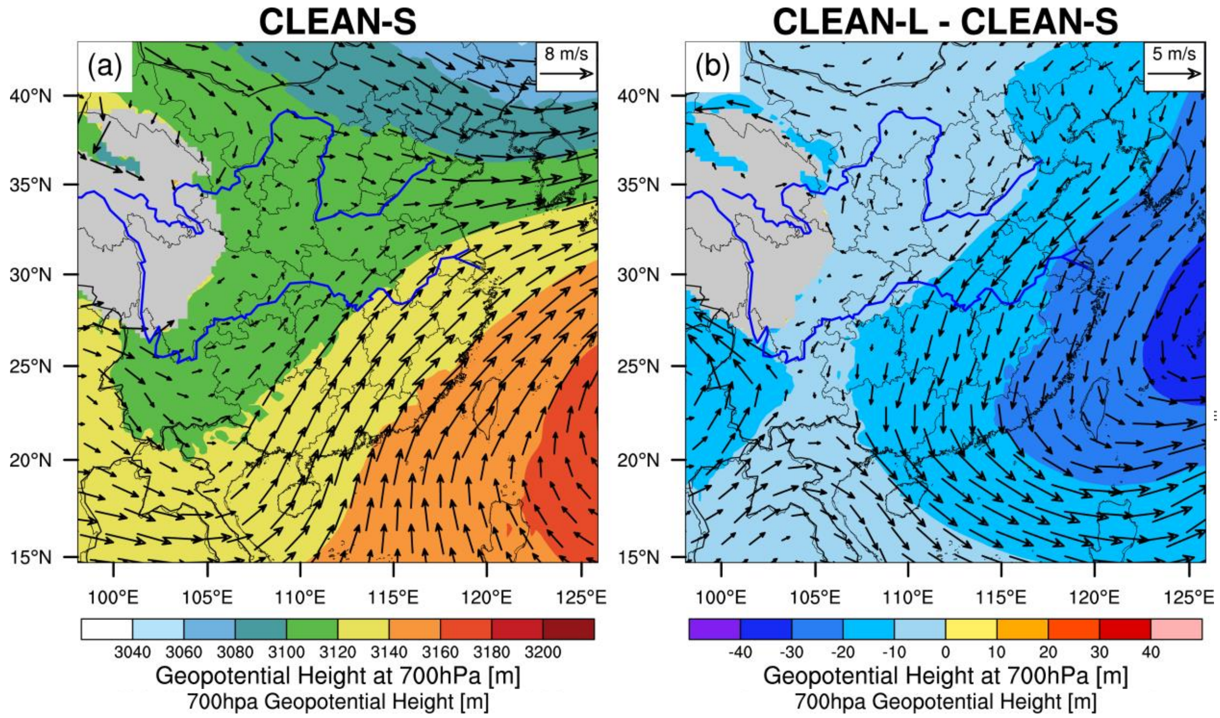
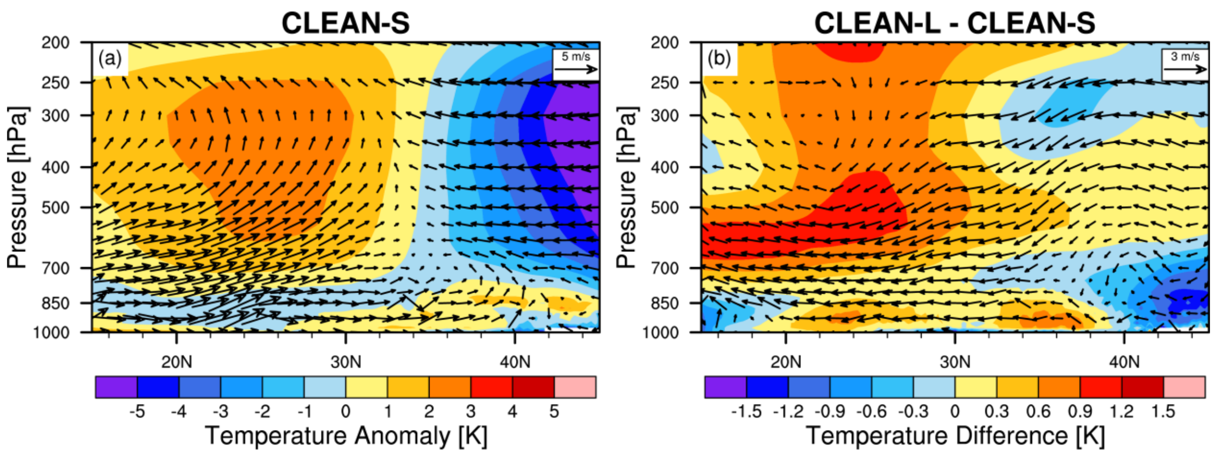


Figure 3. Spatial ~~distribution~~distributions of mean geopotential height and wind fields at 700 hPa ~~Geopotential Height and winds~~ of June and July 2017 from (a) ERA5, (b) CLEAN-S, and ~~the (b~~(c) difference between CLEAN-L and CLEAN-S.



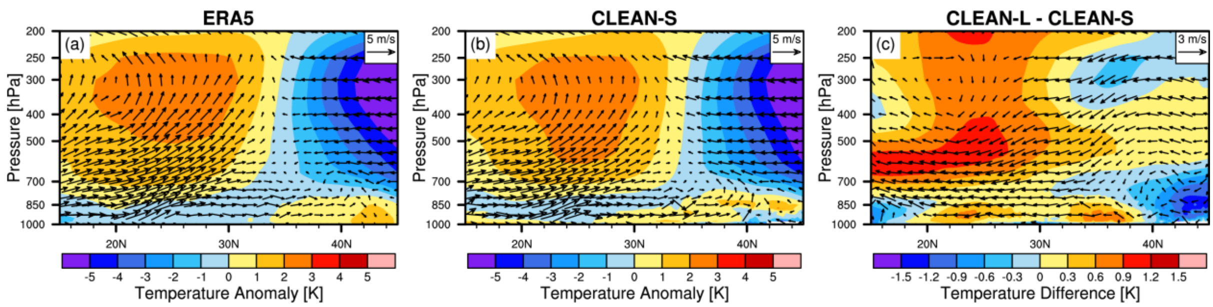
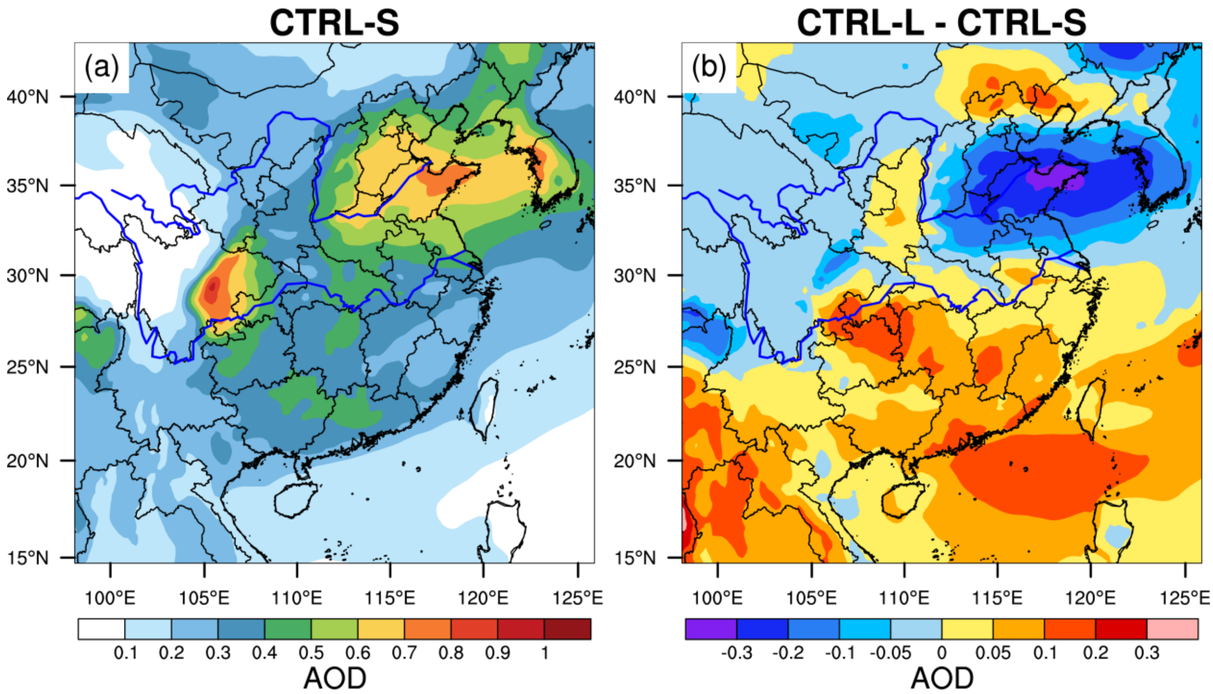


Figure 4. (a, b) The cross-section of meridional temperature anomalies and wind averaged for 105°E and 122°E from ~~(a)~~ the ERA5 reanalysis and the CLEAN-S simulation during June to July, and (~~b~~c) the difference of temperature (not meridional temperature anomalies) between CLEAN-L and CLEAN-S. The meridional temperature anomalies are calculated by subtracting the mean temperature in this latitude range at each pressure level.


CTRL-S
(a)
CTRL-L - CTRL-S
(b)
40°N
35°N
30°N
25°N
20°N
15°N
100°E
105°E
110°E
115°E
120°E
125°E
0.1 0.2 0.3 0.4 0.5 0.6 0.7 0.8 0.9 1
AOD
-0.3 -0.2 -0.1 -0.05 0 0.05 0.1 0.2 0.3
AOD

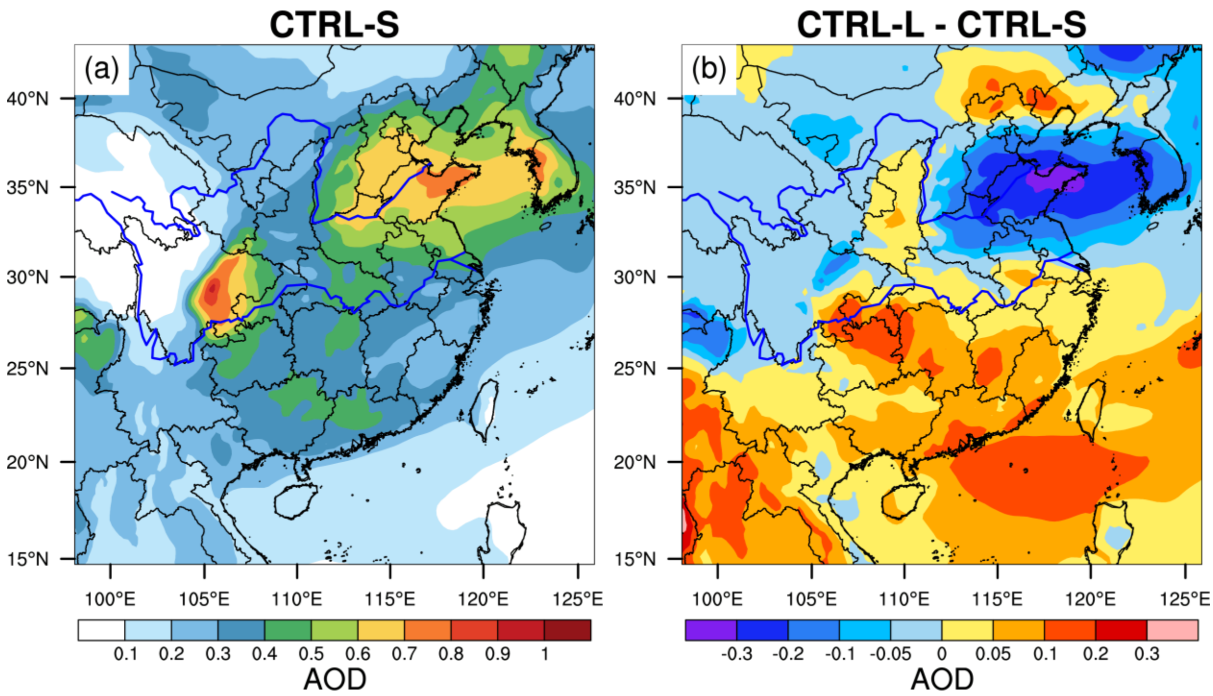
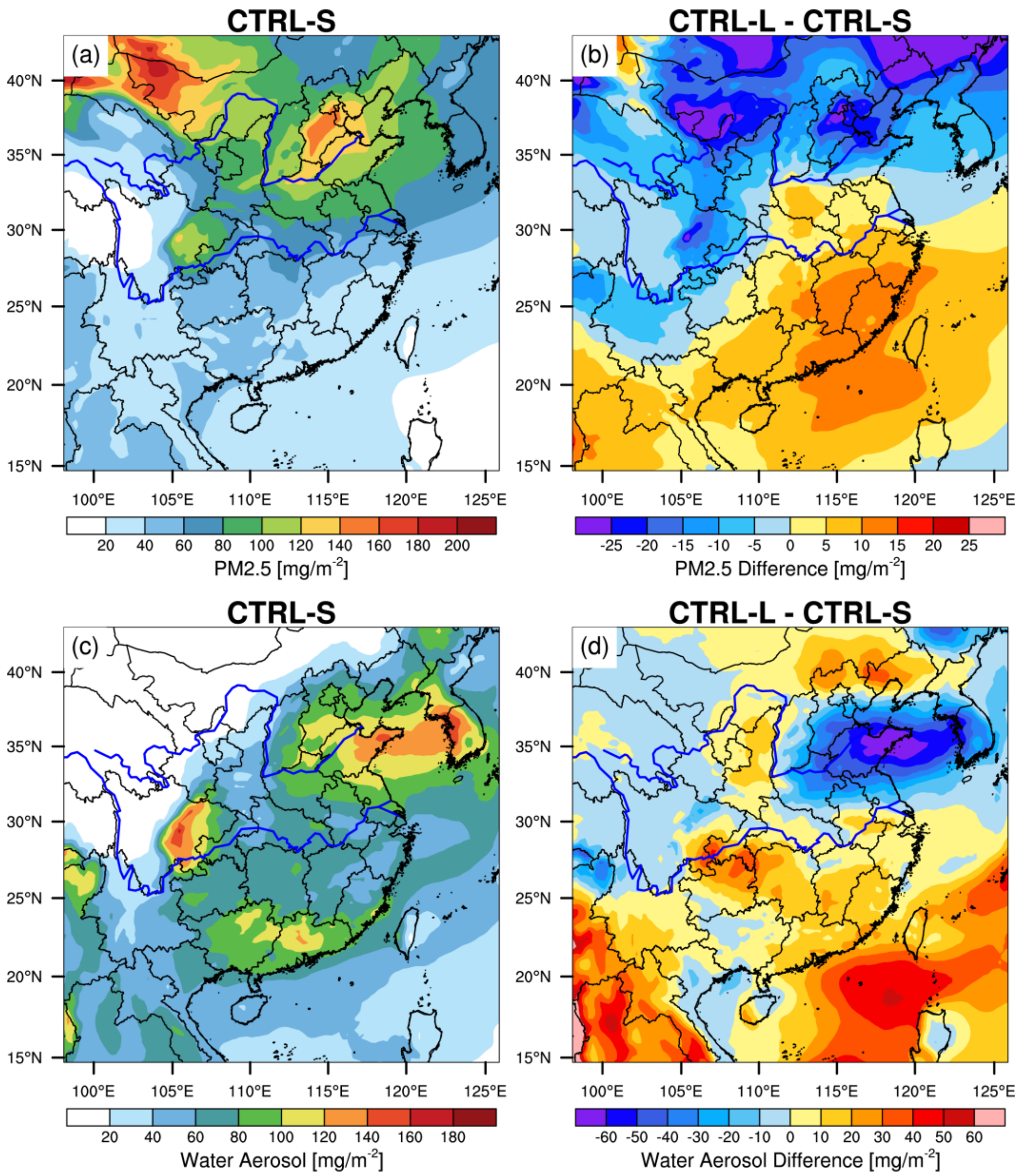


Figure 5. The spatial ~~distribution~~distributions of AOD for June and July of 2017 from the CTRL-S simulation, and the difference between CTRL-L and CTRL-S.


CTRL-S
(a)
CTRL-L - CTRL-S
(b)
PM2.5 [mg/m^{-2}]
PM2.5 Difference [mg/m^{-2}]
CTRL-S
(c)
CTRL-L - CTRL-S
(d)
Water Aerosol [mg/m^{-2}]
Water Aerosol Difference [mg/m^{-2}]
15°N 20°N 25°N 30°N 35°N 40°N
100°E 105°E 110°E 115°E 120°E 125°E

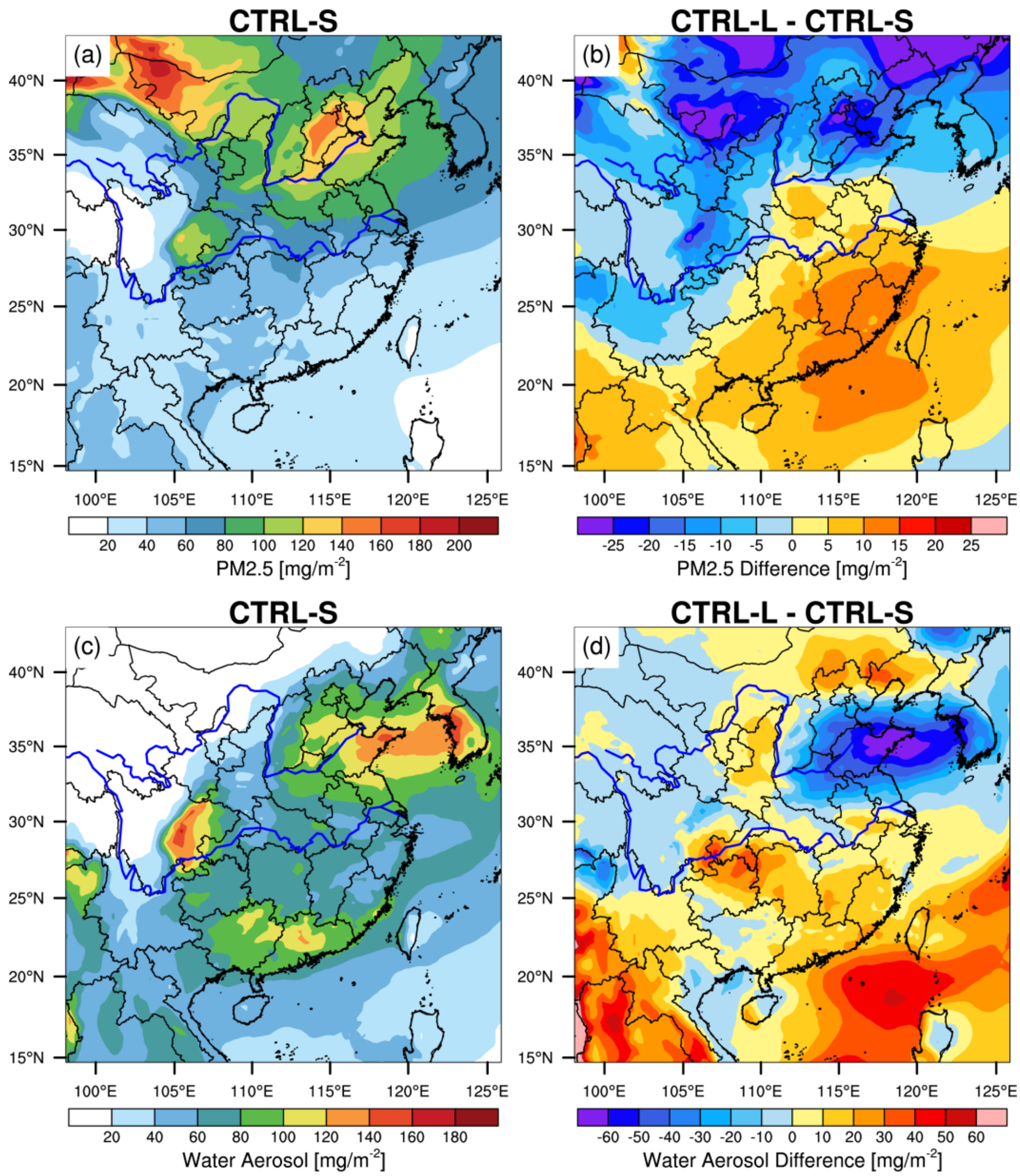
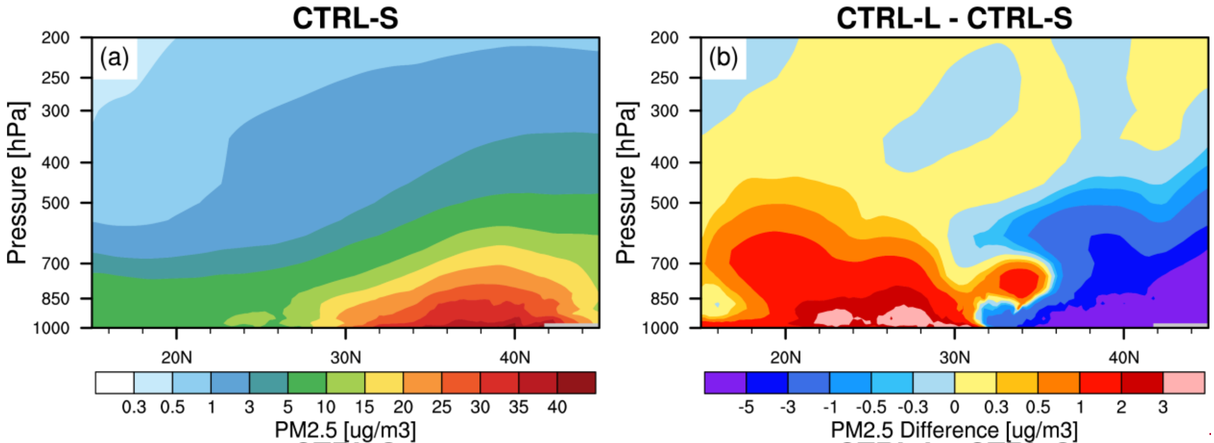


Figure 6. The spatial distributions of column integrated total (a) $PM_{2.5}$ concentration and (c) water content in aerosol averaged for June and July of 2017 from the CTRL-S simulation, and (b~~)(~~, d) the difference between CTRL-L and CTRL-S.



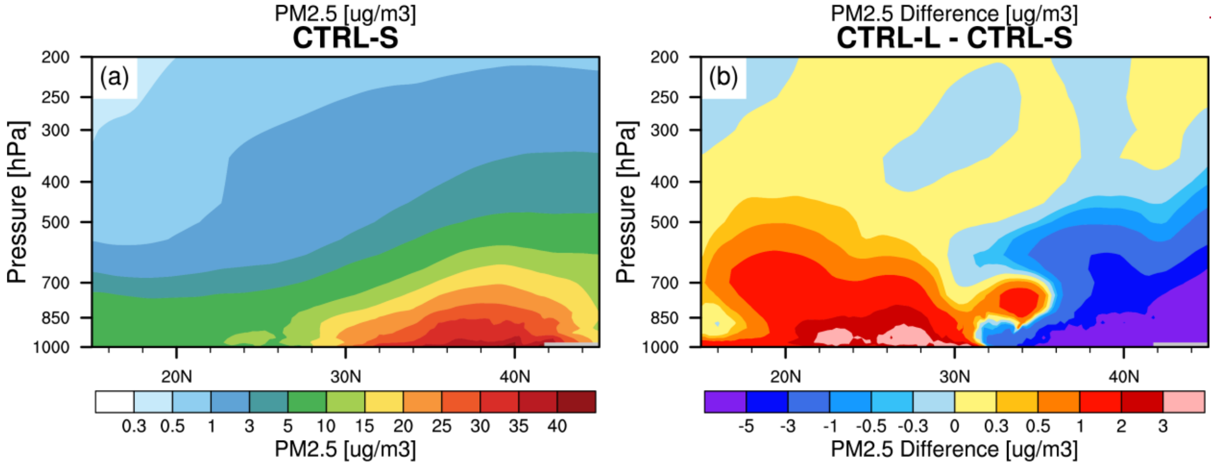
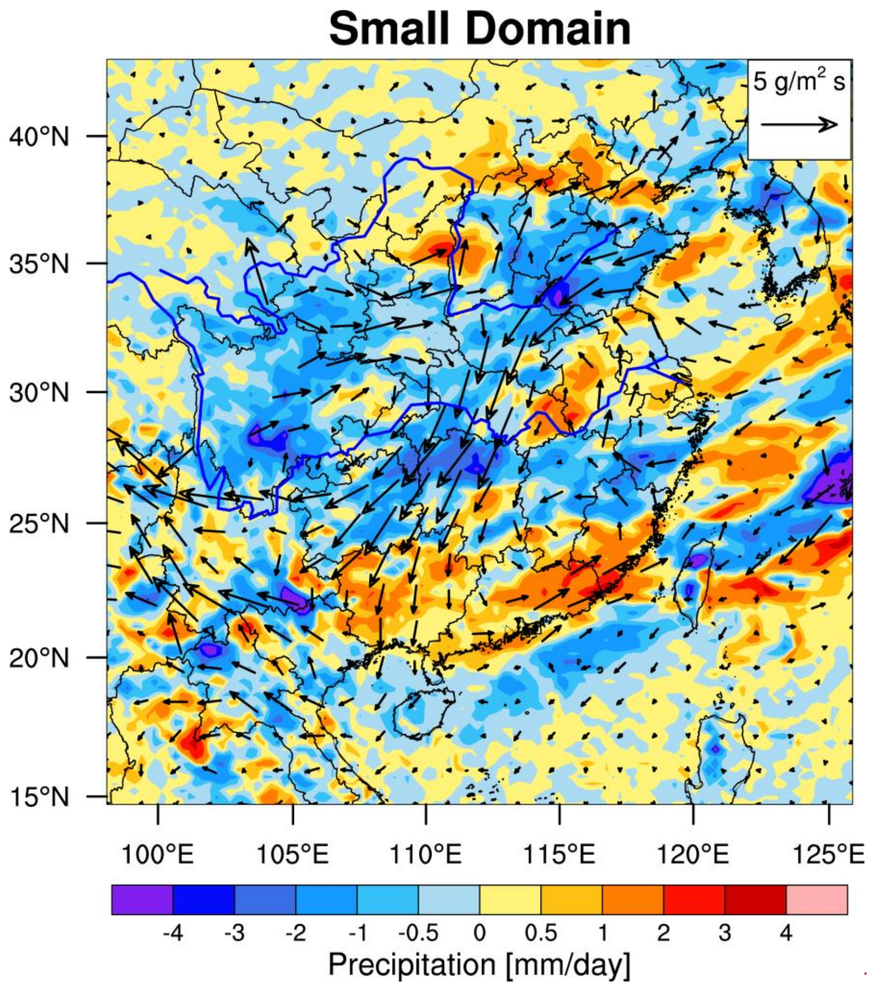


Figure 7. The latitude-height cross-section of (a) total PM$_{2.5}$ averaged between 105°E and 122°E for June and July of 2017 from the CTRL-S experiment, and (b) the difference between CTRL-L and CTRL-S.


Small Domain
5 g/m² s
40°N
35°N
30°N
25°N
20°N
15°N
100°E
105°E
110°E
115°E
120°E
125°E
-4
-3
-2
-1
-0.5
0
0.5
1
2
3
4
Precipitation [mm/day]

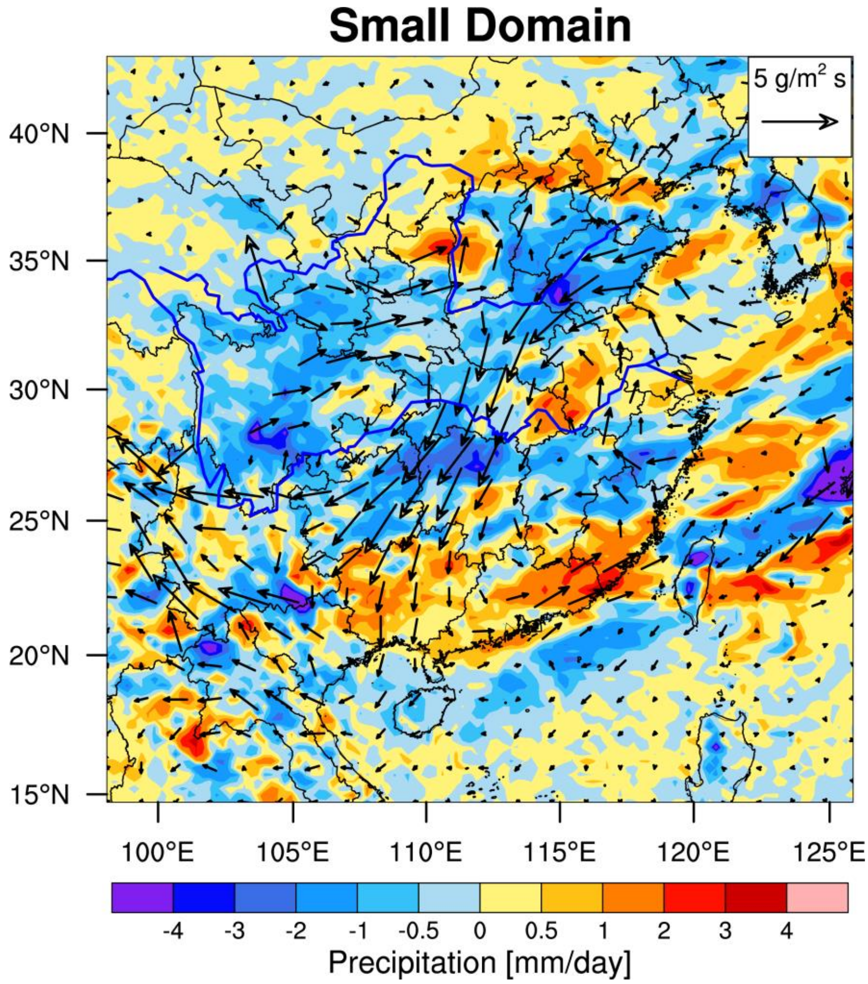
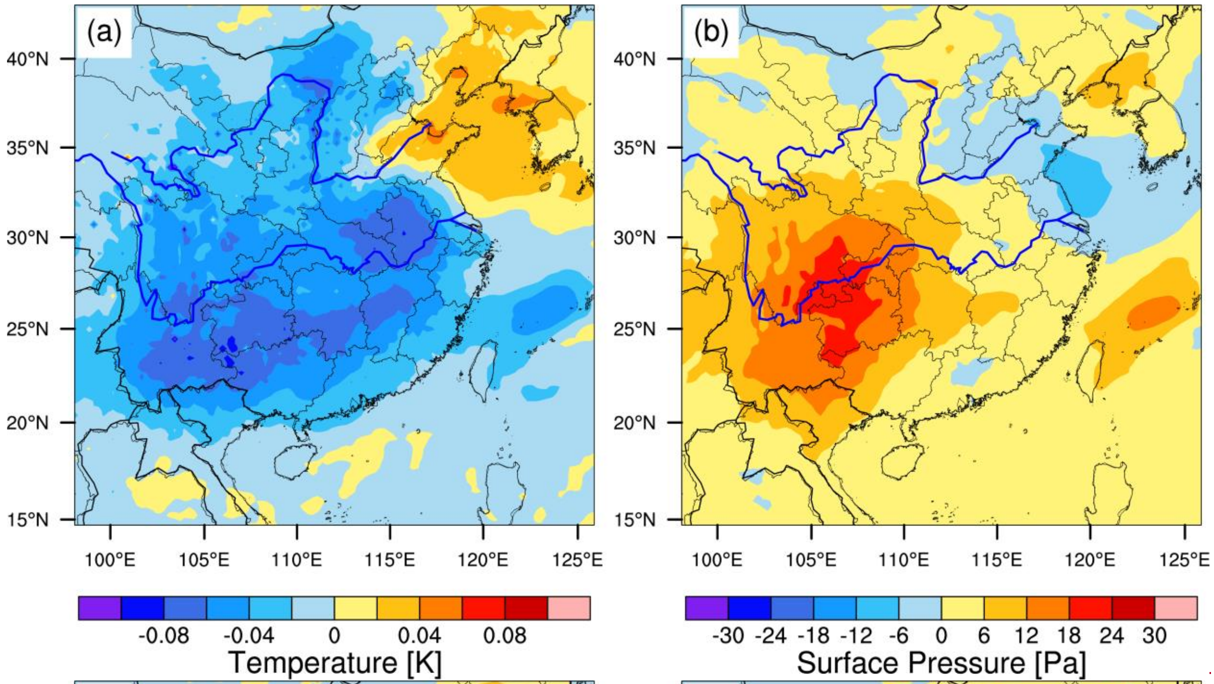


Figure 8. The spatial distributions of aerosol-induced difference (CTRL-CLEAN) of precipitation and moisture transport at 700 hPa averaged for June and July of 2017 from the small domain simulations.



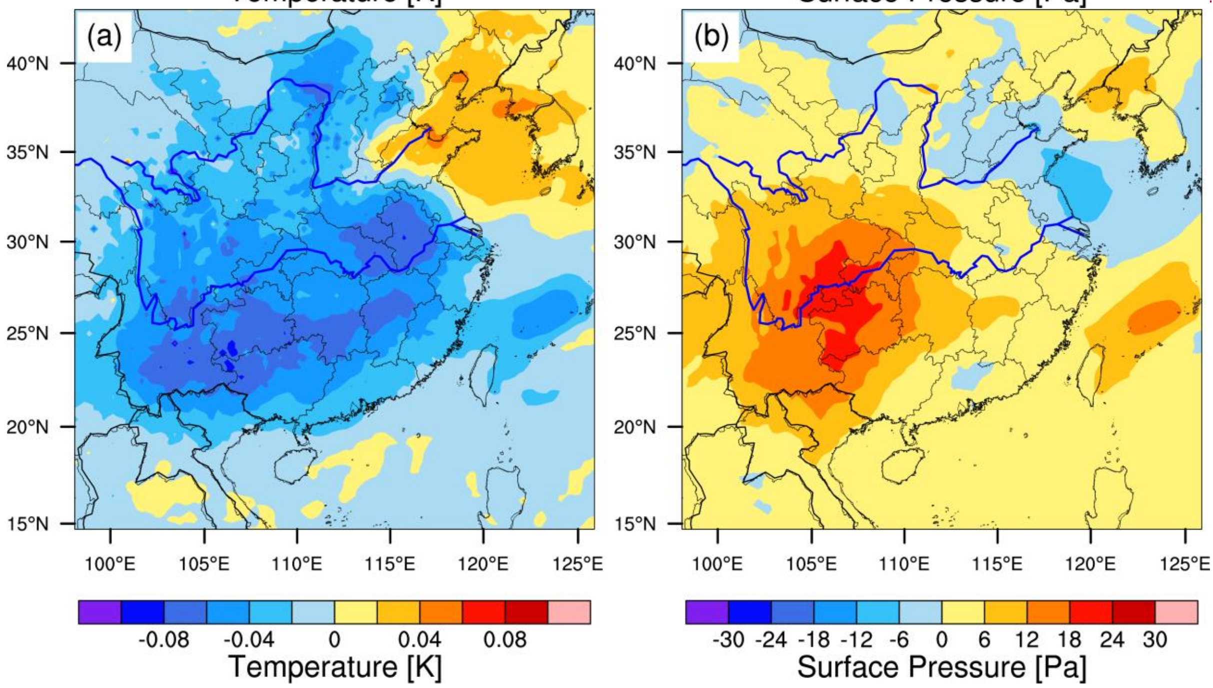
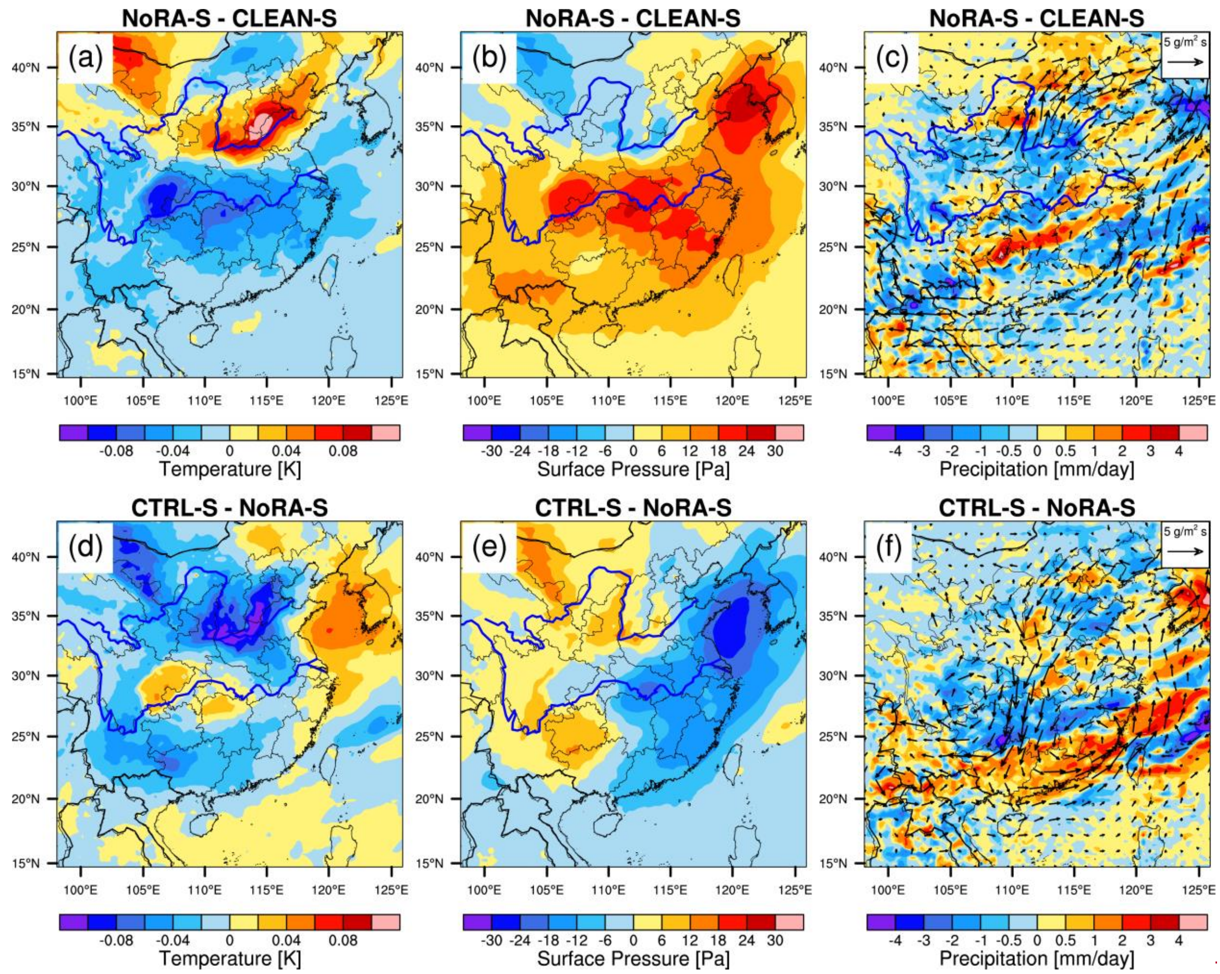


Figure 9. The spatial distributions of aerosol-induced difference (CTRL-CLEAN) of (a) atmosphere temperature below 500 hPa and (b) surface pressure averaged for June and July of 2017 from the small domain simulations. ~~We interpolate the atmosphere~~Atmospheric temperature ~~to the isobaric surface~~is weight-averaged by the layer thickness below 500 hPa ~~and get the atmosphere temperature below 500 hPa by weighted average according to the layer height~~.


NoRA-S - CLEAN-S
(a)
(b)
(c)
5 g/m² s
Temperature [K]
Surface Pressure [Pa]
Precipitation [mm/day]
CTRL-S - NoRA-S
(d)
(e)
(f)
5 g/m² s
Temperature [K]
Surface Pressure [Pa]
Precipitation [mm/day]
-0.08 -0.04 0 0.04 0.08
-30 -24 -18 -12 -6 0 6 12 18 24 30
-4 -3 -2 -1 -0.5 0 0.5 1 2 3 4
100°E 105°E 110°E 115°E 120°E 125°E
15°N 20°N 25°N 30°N 35°N 40°N

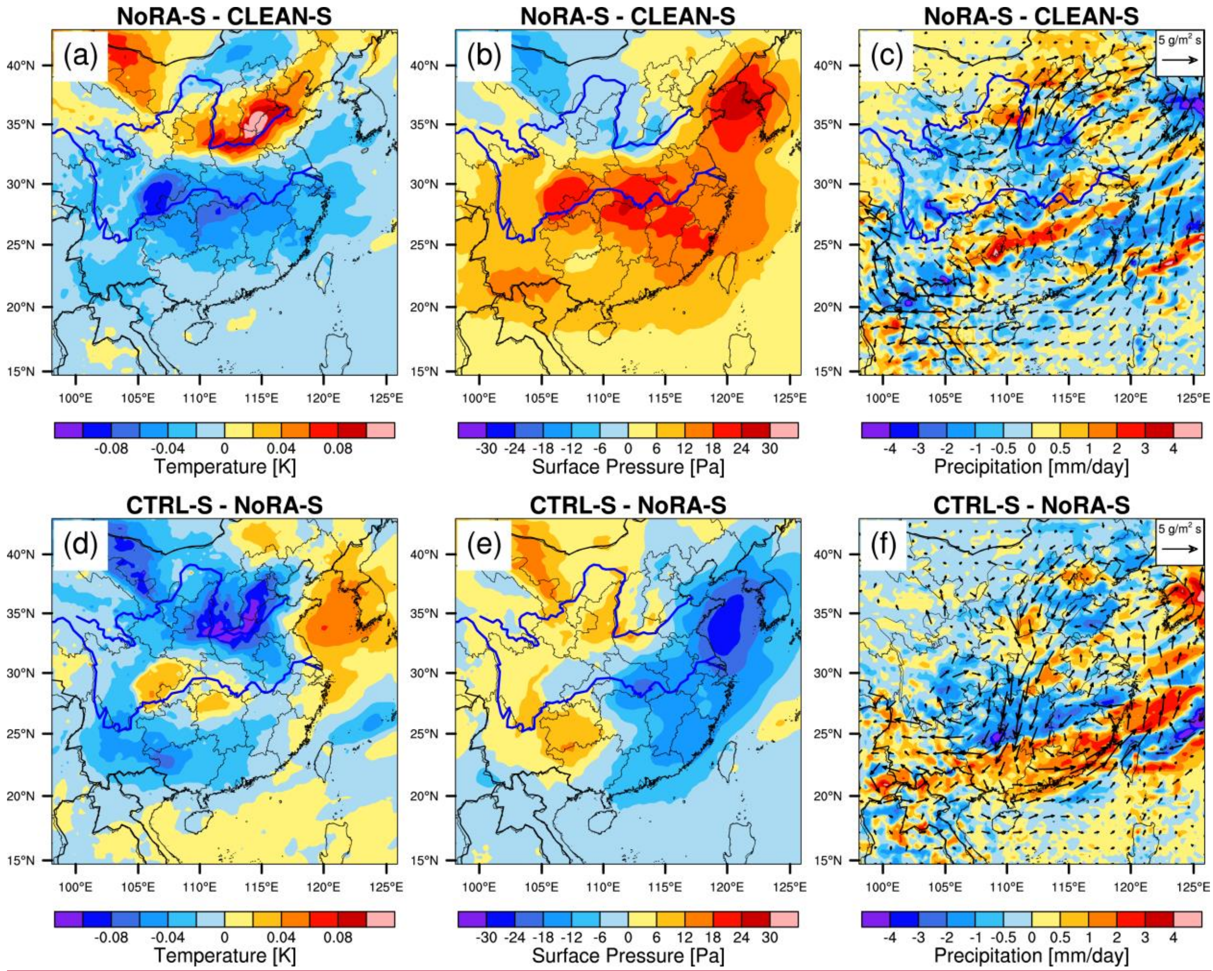
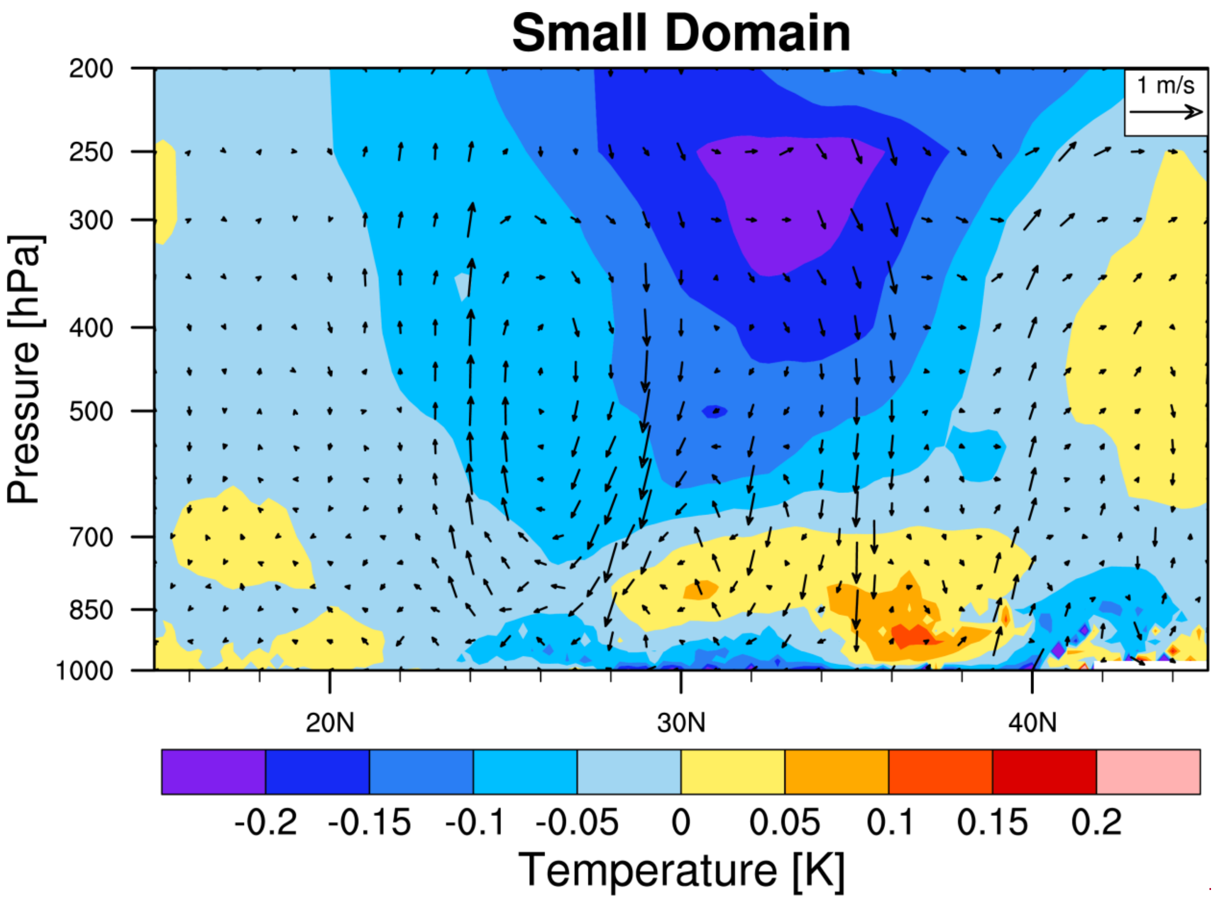


Figure 10. The spatial distributions of Aerosol-Cloud interactions induced difference of (a) atmosphere temperature below 500 hPa, (b) surface pressure and (c) precipitation and moisture transport at 700 hPa averaged for June and July of 2017 from the small domain simulations. And the spatial distributions of Aerosol-Radiation interactions induced difference of (d) atmosphere temperature below 500 hPa, (e) surface pressure and (f) precipitation and moisture transport at 700 hPa averaged for June and July of 2017 from the small domain simulations.


Small Domain
1 m/s
200
250
300
400
500
700
850
1000
Pressure [hPa]
20N
30N
40N
-0.2
-0.15
-0.1
-0.05
0
0.05
0.1
0.15
0.2
Temperature [K]

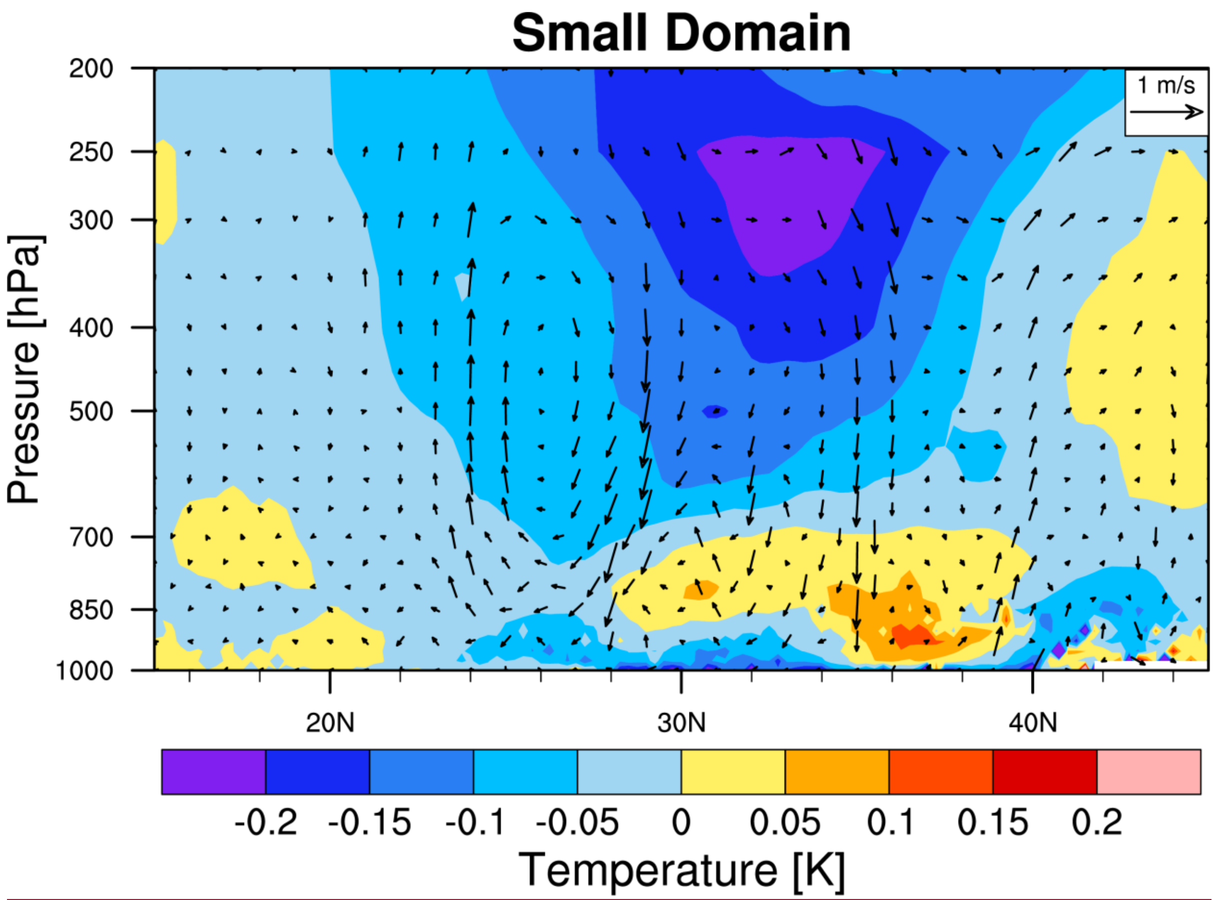
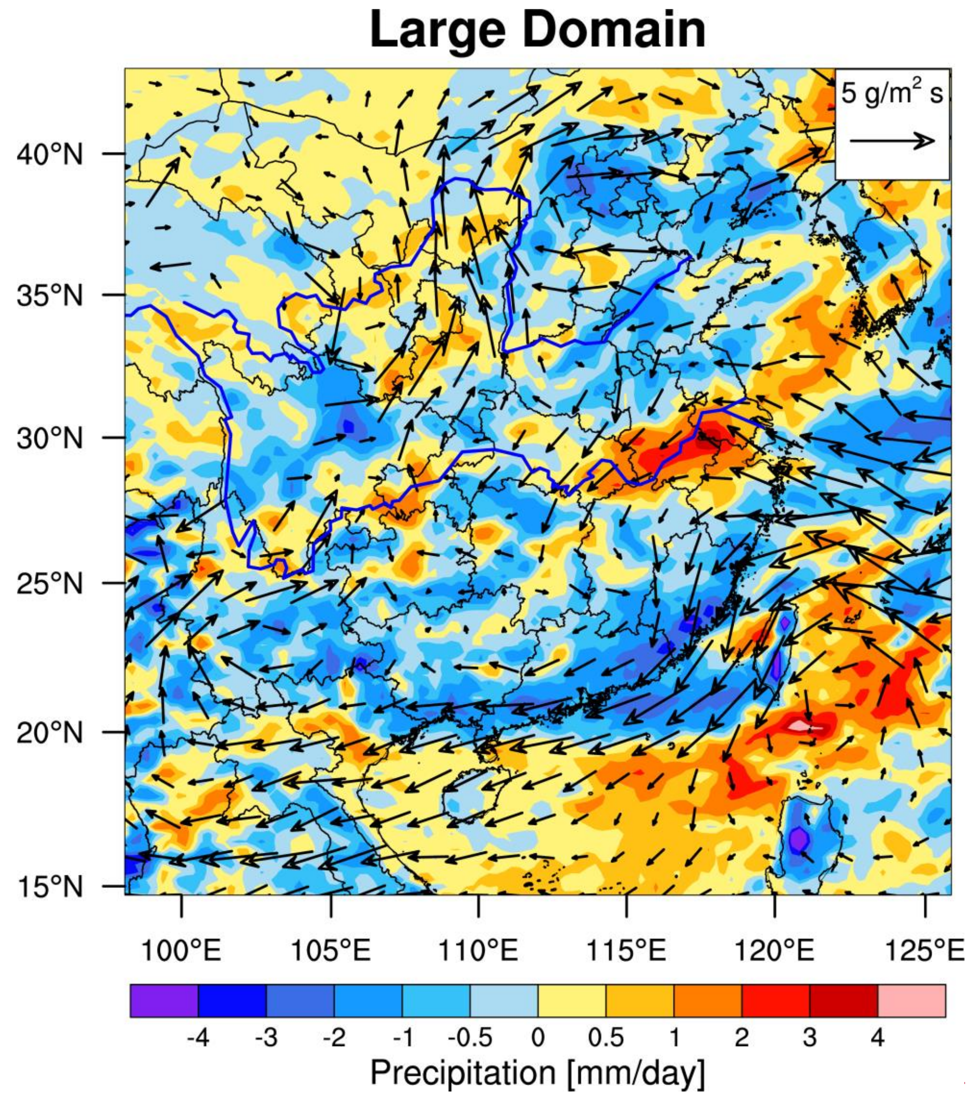

Small Domain
1 m/s
200
250
300
400
500
700
850
1000
Pressure [hPa]
20N
30N
40N
-0.2
-0.15
-0.1
-0.05
0
0.05
0.1
0.15
0.2
Temperature [K]

Figure 11. The latitude-pressure cross-section of aerosol-induced difference (CTRL-CLEAN) of temperature and wind averaged between 105°E and 122°E for June and July of 2017 from the small domain simulation.


Large Domain
5 g/m² s
40°N
35°N
30°N
25°N
20°N
15°N
100°E
105°E
110°E
115°E
120°E
125°E
-4
-3
-2
-1
-0.5
0
0.5
1
2
3
4
Precipitation [mm/day]

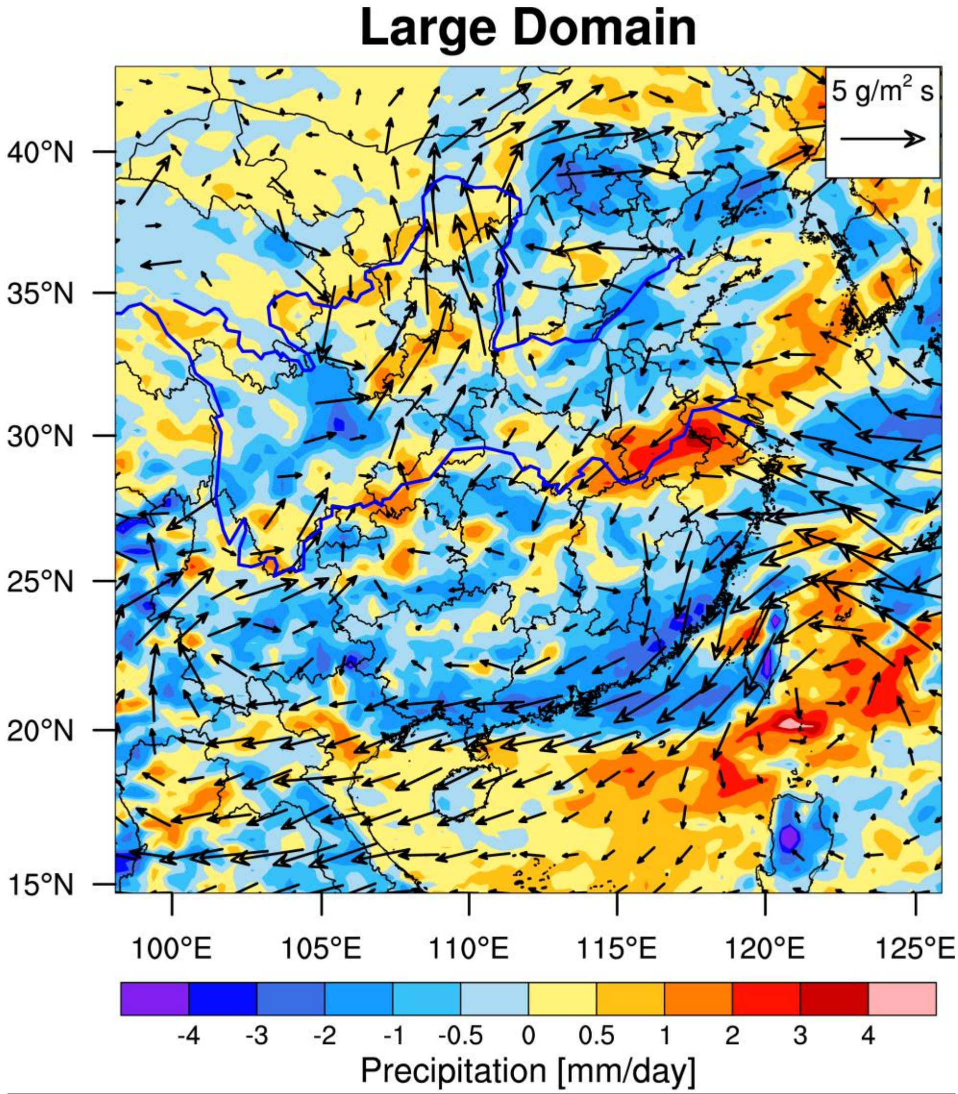
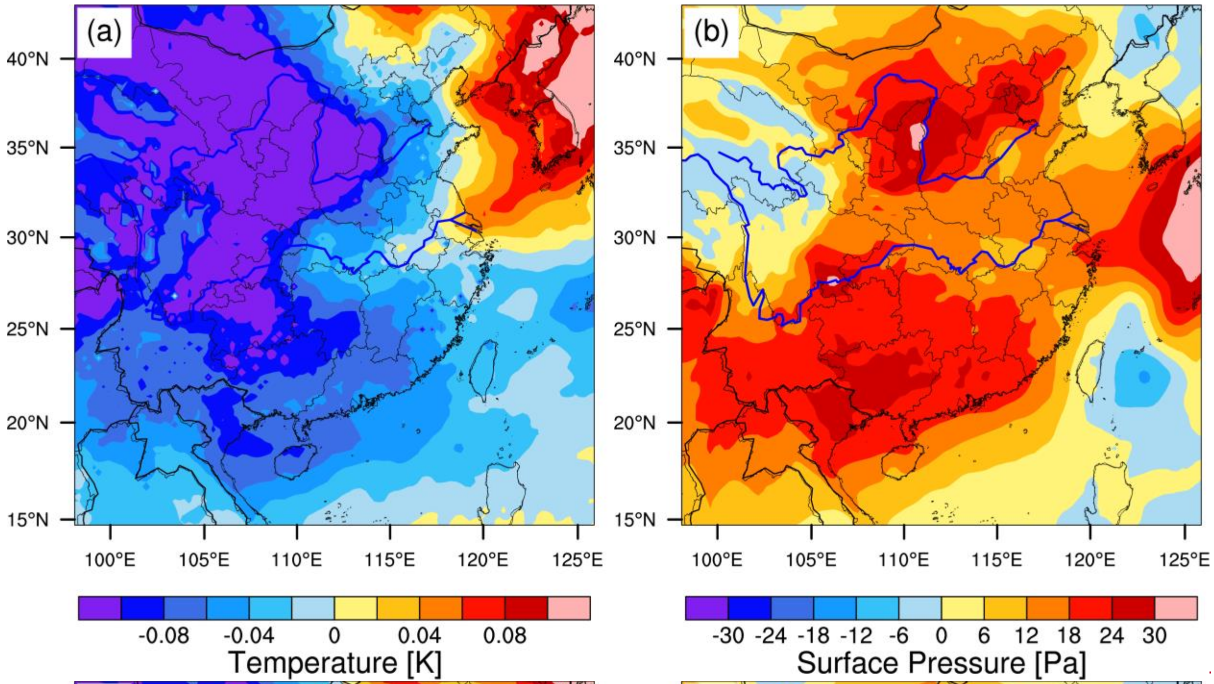


Figure 12. The same as figure 8, but from the large domain simulation.



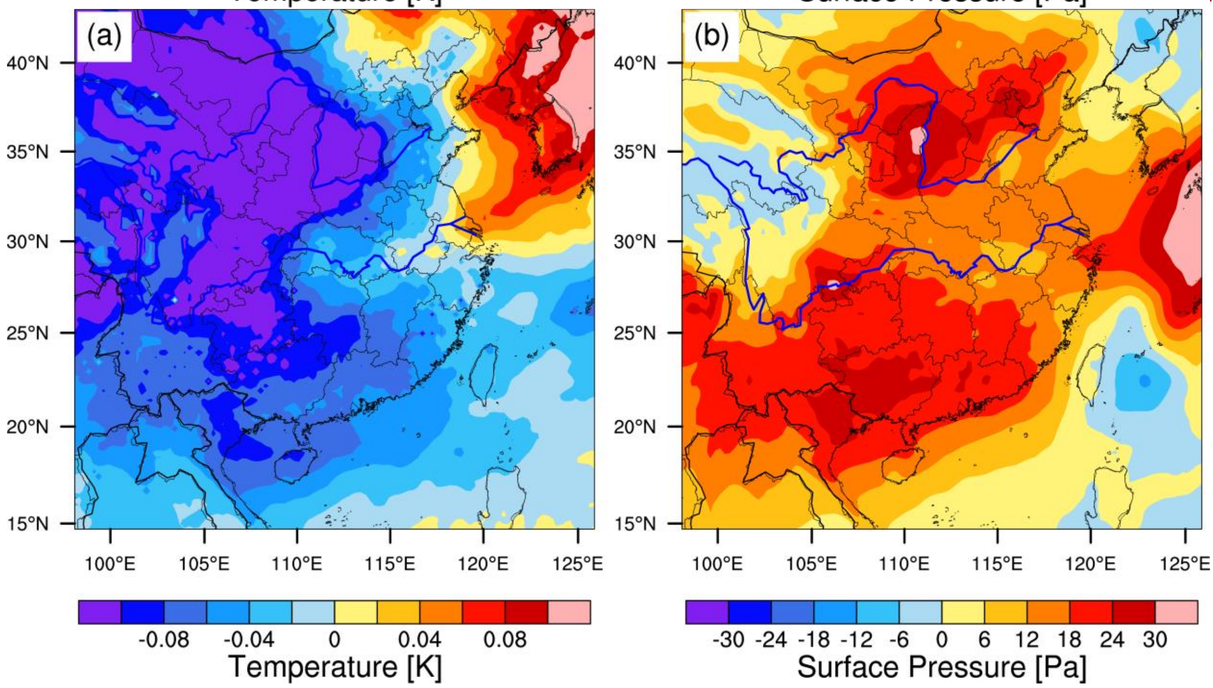
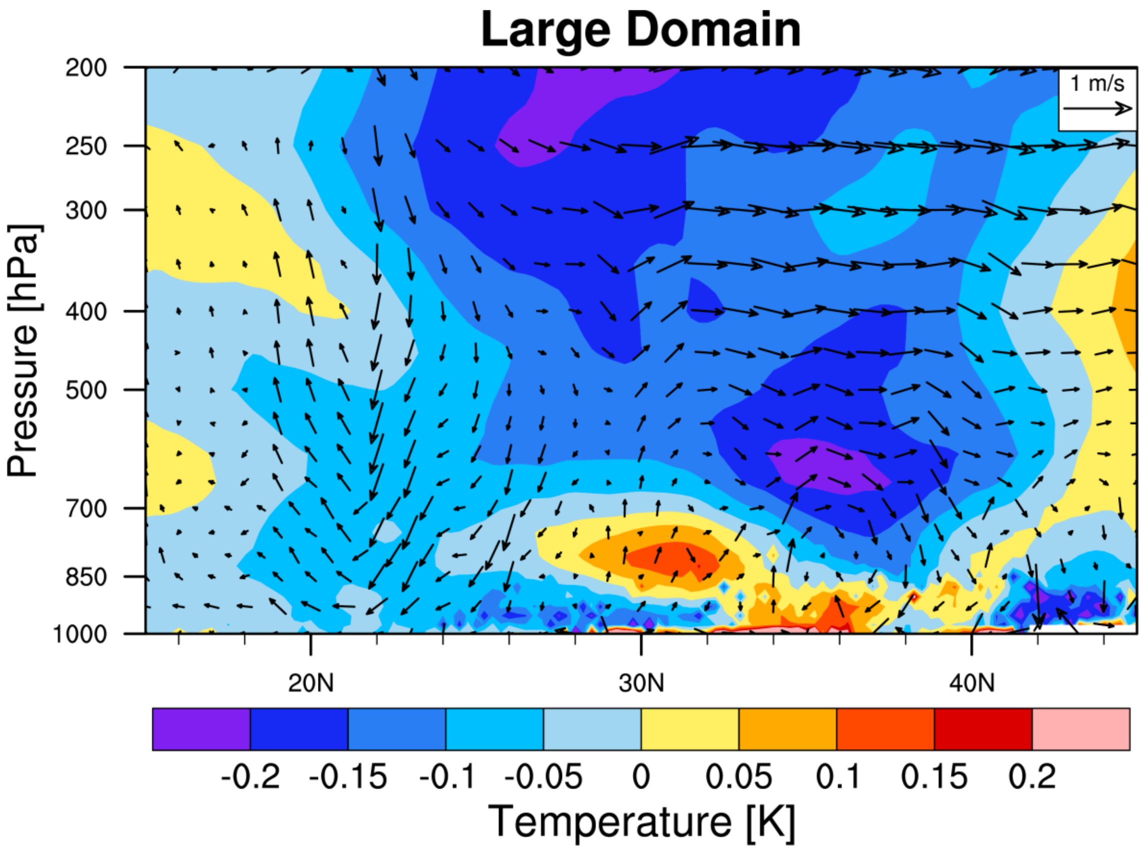


Figure 13. Same as Fig. 9, but from the large domain simulation.


Large Domain
1 m/s
Pressure [hPa]
200
250
300
400
500
700
850
1000
20N
30N
40N
-0.2 -0.15 -0.1 -0.05 0 0.05 0.1 0.15 0.2
Temperature [K]

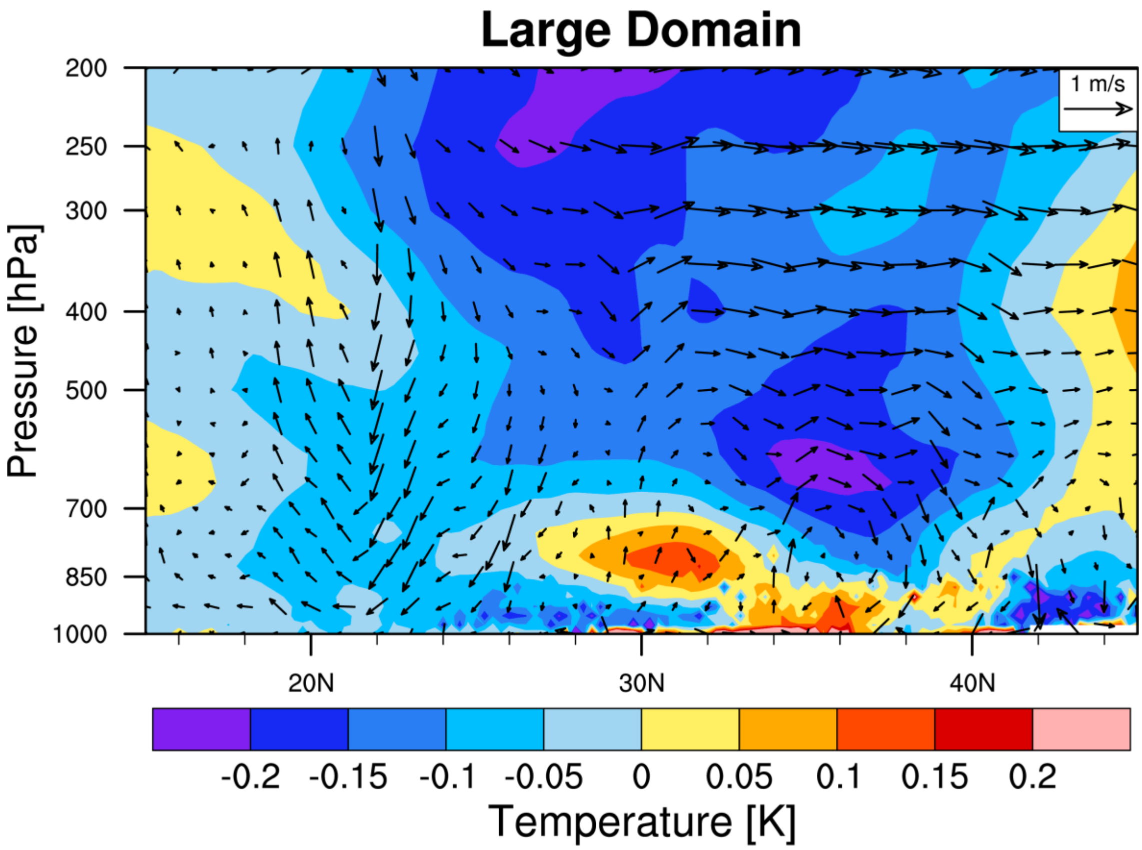
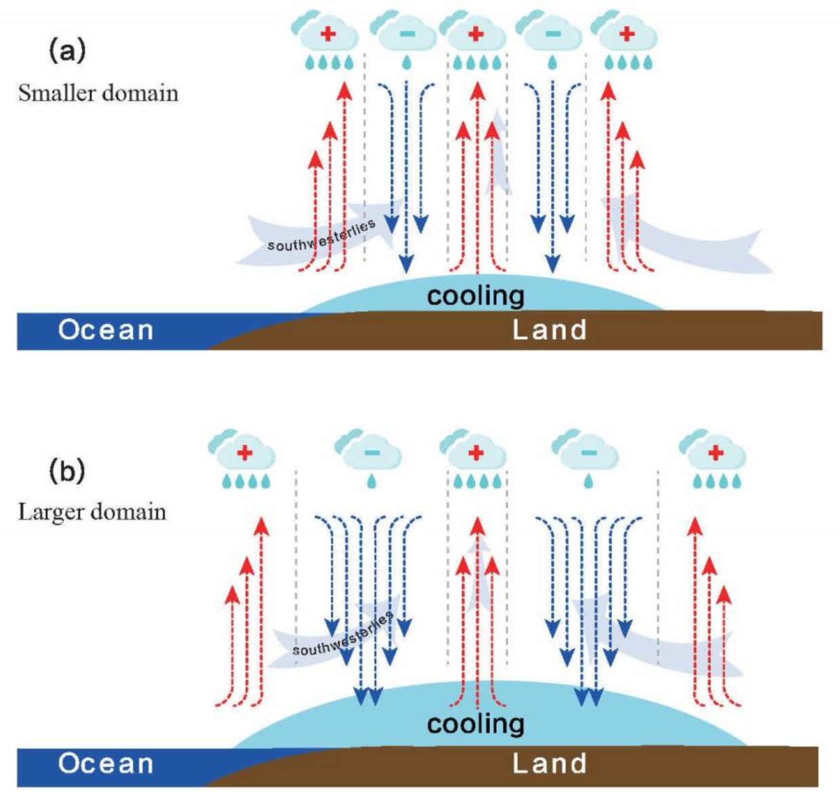


Figure 14. Same as figure 11, but from the large domain simulation.



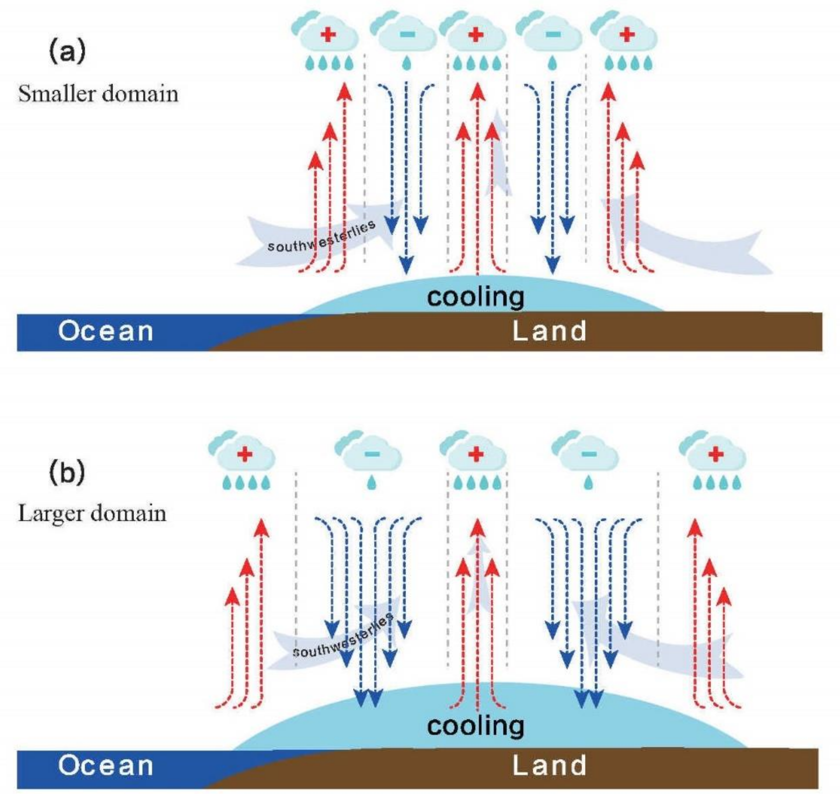


Figure 15. The schematic plot of aerosol ~~impacts~~impact in (a) small domain simulation and (b) large domain simulation over East Asia. The light blue shadow area represents the extent of aerosol induced decrease of lower tropospheric temperature and increase of surface pressure.

The red (blue) vector dash lines represent updraft (downdraft) anomalies. The “+” (“-”) above the region indicates the aerosol-induced increase (decrease) of precipitation.