# A Lagrangian-based Floating Macroalgal Growth and Drift Model (FMGDM v1.0): application in to the Yellow Sea green tides of the Yellow Sea

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Abstract. Massive floating macroalgal blooms in the ocean have haresult indpose manyan array of ecological consequences. T; thus, tracking their drifting pattern and predicting their biomass are important essential for their effective marine the Floating Macroalgal Growth and Drift Model (FMGDMM v1.0), was developed. Based on the tracking, replication, and extinction of Lagrangian particles, FMGDM is capable of determining determine the dynamic growth and drift pattern of floating macroalgal, based on the tracking, replication, and extinction of Lagrangian parti, with clesthe The position, velocity, quantity, and represented biomass of particles being updated are synchronously updated synchronously between the tracking module and the ecological modules. The particle tracking The former is driven by ocean flows and sea surface wind, and the ecological process, while the latter is controlled by the temperature, irradiationsalinity, and nutrients irradiation. -The flow and turbulence fields were provided by Based on the hydrodynamic models of the unstructured grid Finite-Volume Community Ocean Model (FVCOM), and biological parameters -were specified based on parameterized using a culture experiment of Ulva prolifera, a phytoplankton species which causinged the largest worldwide bloom worldwide of the green tide in the Yellow Sea, China, -The FMGDM was applied to simulate, this model was applied to simulate the green tides around the Yellow Sea in 2014 and 2015. The model results, e.g., the simulation result, distribution, and biomass of the green tides, was validated using remote sensing observation data and reasonably modeled the entire process of were validated using the remote sensing observation data. an Given the prescribed spatial initialization from remote sensing observations, the model dereasonably was robust to reproduce modeled the spatial and temporal developments of the entire green tide bloom and its extinction from early spring to late summer. Given the prescribed spatial initialization from remote sensing observa, with tion, the model could provide an accurate prediction for short term (7–8 days) predictions of the spatial and temporal developments of the green tide. With the support of the hydrodynamic model and biological data of macroalgae, this model can forecast floating macroalgae blooms in other regio With the support of the hydrodynamic model and biological macroalgae data, FMGDM nsThis model can serve as a model tool to forecast floating macroalgae blooms in other regions. with the support of the hydrodynamic model and biological data of macroalgae.

#### 1 Introduction

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Floating macroalgae, primarily brown algae and some green algae, extensively occur extensively in oceans. Except for some entirely pelagic species, like *Sargassum*, most floating macroalgae grow in the intertidal zone during their early life stages (Rothäusler et al., 2012). Massive floating macroalgal blooms have frequently recurred in many coastal regions worldwide (Smetacek and Zingone, 2013), and hadcausing deleterious effects on economic activities and ecosystems inof the affected coastal areas (Lyons et al., 2014; Teichberg et al., 2010).

Some floating macroalgae bloomblooms outbreaks-are seasonally, like such as the Sargassum originating from the Gulf of Mexico and the green tide in the Yellow Sea\_(YS), Chinga (YS) (Gower and King, 2011; Liu et al., 2009). Under a suitable temperature and solar radiation environment, the blooms primarily begins in the spring every year. It then is, are advected into the adjacent sea, and growing rapidly in the subsequent floating life stages until they die. The biomass of floating green tide in the YSYS can exceed one million tons in late June (Liu et al., 2013; Song et al., 2015). The A few techniques field and remote sensing observations are have been used to detect the blooming process of floating macroalgae, such as field and remote sensing observations. F However, field samplings, however, observations exhibit site-limitation and are costly, which are difficult to . Moreover, determine ing the overall spatial development of macroalgae blooms in the entire regional sea\_is difficult (Liu et al., 2015). Remote sensing techniques can effectively estimate the coverage and quantify the total biomass (Hu et al., 2019; Wang and Hu, 2016), but they are usually challenging to eannot capture observe the entire development and decay process owing to technical limitations and cloud cover (Keesing et al., 2011). Timely assessment and accurate prediction of coverage and biomass are very essential important for the management and prevention of managing and preventing floating macroalgae bloom.

Numerical simulation is one of the most cost-effective methods to of forecasting spatiotemporal variations of locations and for floating macroalgae. Using Based on athe hydrodynamic numerical numerical numerical hydrodynamichydrodynamich model, we can trace the drift trajectory of floating macroalgae can be determined (Lee et al., 2011; Putman et al., 2018). The biomass, growth, and spatial coverage of the floating macroalgae change dynamically over time, which can be simulated by a The biogeochemical and ecosystem numerical numerical ecosystem models for macroalgae growth are also widely applied in the study of estuaries and coasts (Lovato et al., 2013; Perrot et al., 2014; Sun et al., 2020). The biomass, growth, and spatial coverage of the floating macroalgae change dynamically over time. The growth and mortality are controlled by changing environmental factors, such as temperature, light intensity, salinity, dissolved nutrients, dissolved oxygen, seawater turbidity, and predation by zooplankton (Cui et al., 2015; Shi et al., 2015; Xiao et al., 2016). Incorporating physical drifting models and the biogeochemical growth model appears to be essential to high-precision simulation (Brooks et al., 2018). For efficient management and forecasting, such a coupled physical and ecological model must be capable of predicting spatiotemporal variations of floating locations and biomass. However, the temporal variation of biomass could not be determined via physical modeling; meanwhile, the ecological growth model failed to predict the spatial transportation of floating macroalgae. The efficient management and forecasting of massive floating macroalgal blooms were limited by the lack of high precision coupled physical drift prediction models and ecological dynamics models to predict spatiotemporal variations of floating locations and biomass (Wang et al., 2018).

In this study, we developed a coupled physical ecological coupled growth and drift model of floating macroalgal Floating Macroalgal Growth and Drift Model (FMGDM) was developed for floating macroalgal. This model considers, and the influence of environmental factors, such as temperature, light intensity, and salinity nutrients, on macroalgal growth and depletion. was considered for the ecological module. Driving by the flow and turbulence fields output from the Finite-Volume Community Ocean Model (FVCOM) and parameterized with Based on the regional ocean numerical model system and

sufficient physiological data, <u>FMGDM</u> was used to simulate the drift and growth process of different floating macroalgae blooms can be determined and predicted using FMGDM. Based on the Finite Volume Community Ocean Model (FVCOM), this coupled model was applied to the the recurrent green tides green tide in the of YS duringS. By setting up the model based on the data of the physiological and bloom pattern of *U. prolifera* green tide in YS, the blooming process in the summer of of 2014 and 2015, was simulated in this study, T and the model was validated via result, which was compared with satellite-derived data and field-sampled biomass estimation data, and results indicated that the model were was robust.

The rest of this paper is organized as follows. In Section 2, the development of FMGDM-v1.0, data sources, and the numerical research methods of the green tide in YS are described. In Section 3, the this coupled model is applied to the green tides of YS. The role of physical and ecological driving processes factors are is discussed, with a skill assessment of The stills of the drift module are evaluated by comparinge with the particle tracking trajectories through the comparison without droguedrogued drifters and an evaluation of model reality and accuracy in the evolution process of thes, and the full whole processes of green tides bloomed in the YS in 2014 and 2015 are simulated and verified using satellite data. In Section 4, the uncertainties and prospects of FMGDMM development and application are discussed. Major innovations of this model are summarized in Section 5, following by proposed, and the futur improvements of the model codes and dynamics enutlooks of this model are proposed in Section 5.

## 2 Methodology

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#### 2.1 Model framework

The model system for floating macroalgae growth and drift (FMGDM v1.0) consisted of a Lagrangian particle\_tracking module and an ecological module for macroalgae growth and mortality (Fig. 1). The floating drift process is described by the Lagrangian tracking module, which is developed based on the FVCOM v4.3 offline Lagrangian tracking model (http://fvcom.smast.umassd.edu/) (Chen et al., 2012; Chen et al., 2013; Chen et al., 2021), and driven by surface wind and ocean flows. By contrast, in the macroalgae ecological module of macroalgae, the processes of dynamic growth and mortality process in the floating state are exhibited by particle replication and disappearance, and either the daily growth or mortality rate of each simulated particle is dynamically determined dynamically by the temperature, irradiationsalinity, and nutrients irradiation where the floating particle is in space and time. The position, velocity, quantity, and represented biomass of particles are synchronously updated synchronously between the two modules. All the forcing fields The physical and environmental forcing factors are updated from the regional and local weather, and ocean numerical model system, and marine atlas datasets. Based on the updated locations and biomasses update of simulated particles in spatial and biomass, the coupled individual-based tracking and ecological model, which is applicable in the coverage and biomass simulation of floating macroalgae, is achieved.

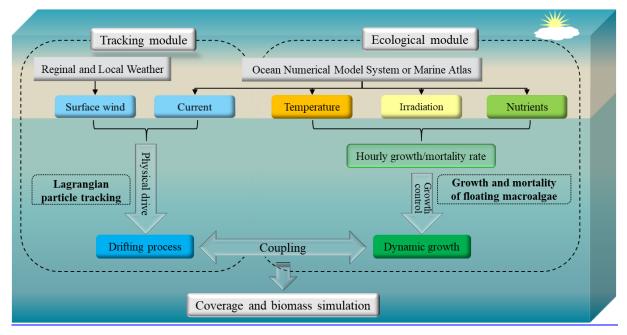


Figure 1. The framework Framework of the physical-ecological coupled model FMGDM v1.0.

## 2.2 Lagrangian particle\_tracking module

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Based on the hydrodynamic model, the Lagrangian particle\_-tracking module was established. The current velocity  $\vec{v}$  is obtained by spatially and temporally interpolating interpolation the three-dimensional (3D) velocity field from the hydrodynamic model. Horizontal and vertical interpolations were carried out via bilinear interpolation, which was also used in the temporal scale. The 10-m-height wind velocity  $\vec{V}_w$ \_Surface wind contributeds to the movement of macroalgae floating aton the sea surface. The windage coefficient  $\kappa$  was determined by Based on the size of macroalgae and the floating depth on the sea, surface wind accounted for 0-2% of wind speed on the drifting of floating macroalgae. Additionally, the wind-induced Stokes drift also accounted for approximately 1.5% of wind on the sea surface. Therefore, a total of 1.5-3.5% ( $\kappa$ ) of the 10 m height wind velocity  $\vec{V}_w$  was considered to determine the additional drifting velocity  $\vec{V}$  of floating macroalgae. Assuming that  $\kappa$  is assumed to be a fixed value, which it does not change with the size of macroalgae in different life stages. The drifting velocity of floating macroalgae patches are is determined using Eq. 1. Additionally, the drift speed is reduced when small patches of macroalgae aggregate as the spatial density  $\rho$  (unit: tons/km²) increases. This effect was regarded as the assemble induced slow moving influence,  $\kappa_d$ , set based on Eq. 2.

$$\vec{V} = \vec{v} + \overrightarrow{V_w} \cdot \kappa - \kappa_d \tag{1}$$

$$\kappa_{d} = \max\left(1 - \frac{p}{75}, 0.1\right)$$
 (2)

To ensure the accuracy of particle trajectory, Eq.  $\frac{3-2}{2}$  is integrated by the fourth-order Runge-Kutta algorithm (Chen et al., 2021), and the time step of calculation  $\Delta t$  is 60 s.

$$X_{t+\Delta t} = X_t + \int_t^{t+\Delta t} \vec{V}(x_t, t) dt$$
 (32)

Dispersion, which was not caused by wind or currents, is also included in the trajectory tracking module. It mainly exhibited a stochastic movement, which was considered horizontal and vertical random walks by adding extra terms to particle trajectory calculation. Since the macroalgae mainly float at the sea surface without significant vertical migration, the vertical random walk wasis not turned on-disabled in the model setting. The horizontal random diffusion of the particles  $\Delta \vec{x_r}$  is also considered in simulation as Eq. 43. The coefficient of horizontal random diffusion,  $K_r$  (r with unit: r m²/s), was set to 200 m²/s, and the time step r Atra for random diffusion was set to 6 s according to Visser's criterion. The unit vector  $\vec{a}$  takes a random direction angle, and the random number r, fits normal distribution, takes a value between 0 and 1.0.

$$\Delta \overrightarrow{x_r}(\Delta t) = \vec{a} \cdot R \sqrt{2K_r \Delta t} \Delta \overrightarrow{x_r}(t)$$

$$= \int_t^{t+\Delta t} \vec{a} \cdot R \sqrt{2K_r t} dt$$
(43)

Therefore, the final position of Lagrangian particle\_-tracking during one-time step  $\Delta t$  can be expressed as:

$$X_{t+\Delta t} = X_t + \int_t^{t+\Delta t} \vec{V}(x_t, t)dt + \Delta \vec{X_r}(t)$$
 (54)

#### 135 **2.3 Ecological module**

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The growth and mortality of macroalgae are controlled by external environmental factors. Surface temperature  $\theta$ , salinity S and irradiation intensity R of surface seawater were used to describe the physiological processes of macroalgae in our model. Daily growth/extinction rate  $\delta$  (% day<sup>-1</sup>) with  $\theta$ , S, and R was determined using laboratory research results and revised according to the actual situation.

$$\delta_{t} = \delta(\theta_{t}, R_{t}, S_{t}, t) \tag{6}$$

The <u>ecological</u> module reflects the <u>process of</u> growth and extinction of macroalgae by the replication and extinction of particles. One initial particle represented a patch with fixed biomass  $(m_0)$  of floating macroalgae, and the value could be adjusted according to needs. It was <u>replicated and randomly</u> released within a 2-km radius of the original location when the represented biomass of the particle exceeded  $2m_0$ . T, and the biomass of the two particles returned to the initial value  $m_0$ . Both particles then <u>undergo underwent</u> drifting and growth/extinction processes independently (Fig. 2a). Additionally, when <u>both</u> two nearby particles had biomass of below  $0.5m_0$ , they were combined to form one particle with <u>a biomass of</u>  $m_0$ , representing the extinction process (Fig. 2b).

The calculation of dynamic change of single particle biomass is expressed as:

$$m_{t+\Delta t} = m_t (1 + \delta_t \Delta t) \tag{7}$$

The total biomass of floating macroalgae throughout the domain can be determined by summing up the biomass of all active particles.

$$M_{\overline{t}} = \sum_{n=1}^{N_{\overline{t}}} m_{\overline{t}, \overline{n}} \tag{8}$$

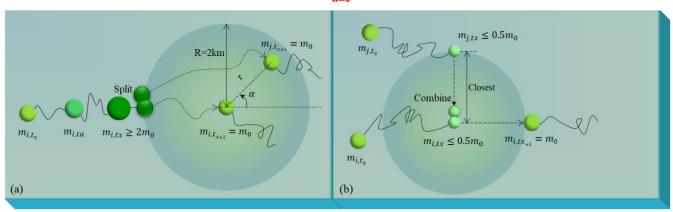


Figure 2. Diagram of the replication (a) and extinction (b) process of simulated macroalgae represented by particles.

The macroalgae growth structure of this ecological module refers to the *Ulva* sp. growth models established by Ren et al. (2014) and Sun et al. (2020). The physiological process of this module is reflected in the absorption and loss of carbon (C), nitrogen (N), and phosphorus (P). Their uptake rates were represented as  $V_c$ ,  $V_N$ , and  $V_P$ , respectively. The loss rates were represented as  $L_c$ ,  $L_N$ , and  $L_P$ , respectively. The calculation of dynamic change of single-particle was expressed as:

$$\frac{dC}{dt} = V_c - L_c \tag{5}$$

$$\frac{dN}{dt} = V_N - L_N \tag{6}$$

$$\frac{dP}{dt} = V_P - L_P \tag{7}$$

We expressed the biomass evolution as carbon. The fresh weight (FW) could be determined by

$$FW = \frac{C}{K_{TOC}} \tag{8}$$

where  $K_{TOC}$  indicates the conversion ratio between C and FW. Based on the physiological characteristics, the value of the conversion value between C and FW was set as  $\frac{968 \text{ mg}}{\text{mg}}$  (Sun et al., 2020).

The total biomass  $(M_t)$  of floating macroalgae throughout the domain can be determined by summing up the biomass of all active particles.

$$M_t = \sum_{n=1}^{N_t} FW_{t,n} \tag{9}$$

The uptake and loss of C, N, and P are controlled by the photosynthesis, respiration, and mortality processes, proportional to biomass. The absorption of C is dependent on the function of photosynthesis f(I), limited by the functions of temperature  $f_p(T)$ , nutrients f(N), and light attenuation of self-shading by macroalgae  $f(\rho)$ .

$$V_c = f(I)f_p(T)f(N)f(\rho) \cdot C$$
 (10)

where  $f(\rho)$  is the effect of self-shading depends on the type of macroalgae and assembled density  $\rho$  (mol C/m<sup>2</sup>). The photosynthesis function f(I) indicates the relation between photosynthesis and irradiation I (Jassby and Platt, 1976).

$$f(I) = P_{max} \tanh\left(\frac{\alpha I}{P_{max}}\right) \tag{11}$$

where  $P_{max}$  is the maximum photosynthetic rate and  $\alpha$  is the photosynthetic efficiency. The changes in internal nutrients quotas have a significant impact on the physiological processes of macroalgae. The N-quota  $(Q_N)$  and the P-quota  $(Q_P)$  represented N:C and P:C, respectively (Ren et al., 2014). The relationship between nutrient quotas and photosynthesis is referred to Droop (1968).  $Q_{Nmin}$  and  $Q_{Pmin}$  are the minimum quota of N and P, respectively.

$$f(N) = min\left[\frac{Q_N - Q_{Nmin}}{Q_N}, \frac{Q_P - Q_{Pmin}}{Q_P}\right]$$
 (12)

The nutrient uptake rate is controlled by the concentration of N and P, and limited by the functions of temperature  $f_p(T)$  and absorption attenuation caused by macroalgae accumulation  $f(\rho)$ . The functions of nutrients uptake rate are referred to Lehman et al. (1975). The absorption of nutrients by macroalgae mainly considers dissolved inorganic nitrogen (DIN) and dissolved inorganic phosphate (DIP). The uptake rate of DIN and DIP are represented as  $V_{DIN}$  and  $V_{DIP}$ , respectively. They are calculated

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$$V_{DIN} = V_{mDIN} \frac{C_{DIN}}{K_{DIN} + C_{DIN}} \frac{Q_{Nmax} - Q_N}{Q_{Nmax} + Q_{Nmin}} \cdot f_P(T) f(\rho) \cdot C$$
(13)

$$V_{DIP} = V_{mDIP} \frac{C_{DIP}}{K_{DIP} + C_{DIP}} \frac{Q_{Pmax} - Q_{P}}{Q_{Pmax} + Q_{Pmin}} \cdot f_{P}(T) f(\rho) \cdot C$$
(14)

The maximum uptake rate of DIN and DIP are represented as  $V_{mDIN}$  and  $V_{mDIP}$ , respectively. The concentration of DIN and DIP are expressed as  $C_{DIN}$  and  $C_{DIP}$ , respectively. The half-saturation coefficient for DIN and DIP are represented as  $K_{DIN}$  and  $K_{DIP}$ , respectively (Sun et al., 2020).  $Q_{Nmax}$  and  $Q_{Pmax}$  are the maximum quota of N and P, respectively.

The C loss is contributed by respiration and mortality. The C loss of respiration depends on the temperature-related function f(T). The C loss of mortality depends on irradiance-related function f(T), nutrients-related function f(N) and temperature-related function  $f_m(T)$ . The temperature limitation functions,  $f_p(T)$ ,  $f_r(T)$ , and  $f_m(T)$ , corresponds to photosynthesis,

respiration, and mortality processes, respectively, w. Where  $R_d$  is the dark respiration rate. When the temperature is unsuitable for the survival of macroalgae,  $f_r(T)$  keeps to a minimum value indicating that minimal respiration. The mortality process replaces photosynthesis as the dominant under severe temperature and light intensity.

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Similar to the loss of C, the uptake and loss of N and P are controlled by the respiration processes, and they are also proportional to biomass. The loss rate of C, N, and P can be calculated by

$$L_c = R_d f_r(T) \cdot C + f(I) f_m(T) f(N) \cdot C \tag{15}$$

$$L_{DIN} = R_d Q_N f_d(T) \cdot C \tag{16}$$

$$L_{DIP} = R_d Q_P f_d(T) \cdot C \tag{17}$$

It should be noted that there is no interaction between this ecological module and the ocean numerical model system since this model is designed for offline computation, which is driven by the ocean model output of physical and ecosystem simulation.

# 2.4 Study area

The first green tide Since their first bloom in the YS outbroke in 2007. Since then, it, green tides hasve become a recurrent phenomenon overphenomenon—over the past 13–15 years (Keesing et al., 2011; Xiao et al., 2020). The major macroalgal species involved in the green tide has been identified as *Ulva prolifera* (Ding and Luan, 2009; Duan et al., 2011). In contrast with some macroalgae that only bloom in certain areas in—such as coastal lagoons and estuaries, the green tides, which accounting for most trans-regional macroalgal blooms worldwide (Liu et al., 2013), is much more complicated, both in spatial and temporal variations. The *U. prolifera* green tides in the YS primarily originates from the coast of Jiangsu Province, primarily the coast of Yancheng and Nantong, and can drift northward to the southern shore of Shandong Peninsula and the coastal region of the Korean Peninsula (Liu et al., 2013; Son et al., 2012) (Fig. 3). Many loosely floating propagules of *U. prolifera* were provided from mid-Apr to mid-May every year (approximately 4,000–6,000 tons), which could float and grow in the YS (Fan et al., 2015).

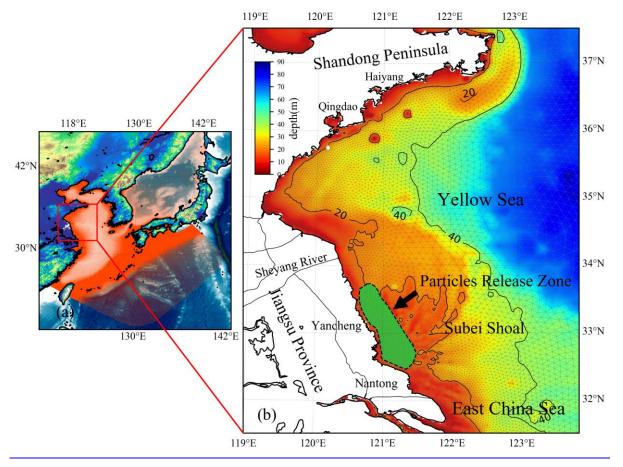


Figure 3. (a) Location of the YS, China. The red mesh indicates the high-resolution triangle grids of ECS-FVCOM. (b) The enlarged view of the area and bathymetry bounded by the solid red line rectangle in panel (a). The green area surrounded by a black dotted line in the Subei Shoal indicates the initial release zone of simulated particles. The green squares, blue bullets, and triangle indicate the positions of particles released on the coast, southern and northern regions of the Subei Shoal, respectively. The lime green stars indicate that the initial position of the tracking simulation only considers the action of ocean flows. Black dots indicate the position of the 47 initial particles released for simulation in 2014 and 2015.

#### 2.5 Data sources

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#### 2.5.1 Surface wind

The wind data at 10 m above sea level derived from the surface wind dataset from the European Centre for Medium-Range Weather Forecasts (ECMWF) are available at: <a href="https://www.ecmwf.int/en/forecasts/datasets/">https://www.ecmwf.int/en/forecasts/datasets/</a>. The wind is interpolated to the triangular grids, covering YS, East China Sea, Bohai Sea, and Japan Sea, in spatial and time scale. The interpolated wind data was were used as surface forcing for ECS-FVCOM, with the-a spatial resolution of 0.125° and temporal resolution of 1 h.

# 2.5.2 Satellite data

The distribution area and density of the green tides can be estimated from satellite data (Hu et al., 2019; Qi et al., 2016). In this study, the spatial distribution and growth density of the green tides green tides's spatial distribution and growth density were validated using satellite data. The Moderate Resolution Imaging Spectroradiometer with Terra sensor (MODIS-TERRA) measured the satellite data of he green tides in the YS in 2014 and 2015, are and the data are available from https://terra.nasa.gov/about/terra-instruments/modis. the Moderate Resolution Imaging Spectroradiometer with Terra sensor (MODIS-TERRA). In addition, the biomass quantified based on the satellite data from Hu et al. (2019) was used to verify the simulated *U. prolifera* biomass. Blocked by clouds, rRemote sensing techniques exhibit difficulty in detecting small patches of floating macroalgae and often-may easily fail to capture the early status of the green tides (Garcia et al., 2013). In this study, oonly a few remote sensing images observations that were not blocked by clouds can\_be used for be used for result verification. The remote\_sensing dataset from https://www.ghrsst.org/, Group for High-Resolution Sea Surface Temperature (GHRSST), is used for data assimilation of sea surface temperature in the model system. The GHRSST dataset is daily based, with a spatial resolution of 0.01 degrees.

#### 2.5.3 Drifting trajectory data

The drifting dataset used to evaluate the skills of the tracking module is composed of two parts: the trajectory data of satellite-tracked surface drifters released from the Subei Shoal in 2012 (Bao et al., 2015), and the subsurface drogued drogue-drifters tracking data in the inner shelf of the ECS in 2017. The surface drifters contained four 40-cm width, 70-cm height rectangular sails, and a large central buoy (Bao et al., 2015). The subsurface drifter was constructed by a 67-cm diameter, 6-m height cylindrical subsurface sail, and a 28-cm diameter central buoy.

#### 2.5.4 Nutrients data

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The seasonal nitrate and phosphate datasets of the YS were are obtained from the 1 degree-resolution World Ocean Atlas 2018 (Garcia et al., 2019), with a spatial resolution of 1 degree, and merged with the datasets from the marine atlas of the YSellow (Wang et al., 1991). With combined two datasets, the nutrients at the sea surface (April to August) were are applied to this simulation through temporal interpolation.

# 2.6 Hydrodynamic model

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An unstructured-grid FVCOMinite Volume Community Ocean Model (FVCOM) adapts to the second-order accurate discrete flux algorithm in anthe integral form to solve the governing equations on an unstructured triangular grid, which provides excellent mass and momentum conservation during the calculation (Chen et al., 2006; Chen et al., 2007; Chen et al., 2003; Ge et al., 2013). To better identify the ocean circulation along with the shelf break and deep ocean, the semi-implicit discretization, which could avoid the adjustment between two-dimensional external mode and three-dimensional internal mode, was applied. With this configuration, the ocean circulation, as well as the astronomical tide around the East China Sea (ECS), YS, and adjacent regions, could be reasonably determined (Chen et al., 2008; Ge et al., 2013). An integrated high-resolution numerical model system for the ECSast China Sea (ECS-FVCOM) based **FVCOM** v4.3 (http://fvcom.smast.umassd.edu/fvcom/) was established and comprehensively validated using observational data (Chen et al., 2008; Ge et al., 2013). The high-resolution triangularle grids of the ECS-FVCOM domain covers the YS, ECSEast China Sea, Bohai Sea, and Japan Sea, which have horizontal resolutions varying from 0.5-1.5 km in the estuary and coastal region, approximately 3 km in the path of the Kuroshio, and 10–15 km along the lateral boundary in the north Pacific region (Fig. 3a). A total of 40 generalized sigma 40 layers are considered in the vertical, including five uniform layers with a thickness of 2 m specified in the sea surface and bottom to resolve better to better resolve surface heating and wind mixing, and bottom boundary (Chen et al., 2008). The ocean bathymetry was retrieved and interpolated from ETOPO1 (https://ngdc.noaa.gov/mgg/global/global.html). The initial temperature/salinity field and the volume transports along the open boundary of ECS-FVCOM were interpolated and retrieved, from HYCOM+NCODA Global 1/12° Analysis data (GLBA0.08), and eight major tide harmonic constituents (M2, S2, K2, N2, K1, O1, P1, and Q1), which are obtained from TPXO 7.2 Global Tidal Solution (Egbert and Erofeeva, 2002), were used along the open boundary (Ge et al., 2013). The freshwater discharge of the Yangtze River and Qiantang River (source: http://www.cjh.com.cn/) was added to the upstream river boundary. Surface wind and radiations from ECMWF were used in ECS-FVCOM as surface forces. In addition, the GHRSSTroup for High-Resolution Sea Surface Temperature (GHRSST) dataset was applied to better determine the sea surface temperature using the model-data assimilation. The simulation time was set from March 29 to September 1, thus covering early spring to late summer. The water velocity, temperature, salinity from ECS-FVCOM were fed into the FMGDM as input variables.

## 2.7 Model settings

The model was parameterized according to the physiological characteristics of the floating *U. prolifera*. Furthermore, 2 % of wind speed was accounted for by the drifting of floating *U. prolifera*, and 1.5 % of wind was accounted for by the wind-induced Stokes drift. Therefore, the coefficient  $\kappa$  was set to 3.5 % (Bao et al., 2015). Based on many physiological research studies on *U. prolifera*, the optimal range of temperatures, irradiance, and salinity can be determined (Table 1). Owing to the protection of the dense branch, floating *U. prolifera* are more tolerant of high irradiance in fields than in the laboratory (Xiao et al., 2016). In the optimal environment range, the daily growth rate of floating *U. prolifera* can reach 10.6–16.7 % (Xiao et al., 2016). Therefore, the relative growth rate of the *U. prolifera* growth module was set refer to the results of physiological research and actual growth in fields. (Table 2).

Table 1. Optimal range of temperatures, irradiance and salinity during U. prolifera growth

Environmental factors	Lower limit	Suitable	Upper limit	Reference
Temperature (°C)	<del>10</del>	<del>14 27</del>	<del>30</del>	Xiao et al. (2016)
<del>Irradiance</del>		<del>200 2400</del>	<del>2400</del>	Cui et al. (2015)
(µmol photons m <sup>-2</sup> -s <sup>-1</sup> )	_	<del>200 - 2/100</del>	<del>2/100</del>	Xiao et al. (2016)
Salinity (psu)	8	<del>26 32</del>	_	Xiao et al. (2016)

Table 2. Relative growth rate setting of *U. prolifera* growth model at various temperatures, irradiance and salinity.

Environmental factors			Deletive energth note (0/ d=1)
Salinity (psu)	<del>Temperature (°C)</del>	Irradiance (µmol photons m <sup>-2</sup> s <sup>-1</sup> )	Relative growth rate (% a 1)

	< <u>-10</u>	<u>≥ 0</u>	0	
	<del>10 14</del>	<del>0 200</del>	<del>2</del>	
		<del>200 3200</del>	7	
<del>26 32</del>		<del>≥ 3200</del>	2	
	<del>14 25</del>	<del>0 200</del>	<del>5</del>	
		<del>200 3200</del>	<del>22</del>	
		<del>≥ 3200</del>	<del>5</del>	
		<del>0 200</del>	0	
	<del>25 27</del>	<del>200 3200</del>	<del>2</del>	
		<del>≥ 3200</del>	<del>_2</del>	
	<del>27 29</del>	<del>≥ 0</del>	<del>-12</del>	
	<del>≥ 29</del>	<del>≥ 0</del> ≥ 0	<del>-90</del>	

To determine the drift trajectory of *U. prolifera* in the natural state and understand the contribution of these physical factors, two particle tracking experiments were designed. Three group particles were released on May 1, 2014 in the coastal, southern, and northern regions of the Subei Shoal respectively (Fig. 3b). The roles of ocean circulation and surface wind were considered in this tracking experiment, and the simulation lasted 120 d. Meanwhile, the other tracking experiment was configured with mostly the same settings, but the effect of surface wind on floating particles was disregarded. In the second experiment, initial particles were also released on May 1, 2014 but only around the Subei Shoal (Fig. 3b).

Most importantly, two realistic dynamic growth simulations were conducted. To verify the general applicability of the model, the growth and drift processes of *U. prolifera* in YS in 2014 and 2015 were simulated, respectively, with identical model configurations. In the two simulations, a total of 47 particles were initially released into the Subei coast on May 1, 2014 and 2015, uniformly distributed in space (Fig. 3b). One initial particle represented a patch of 100 tons of floating *U. prolifera*. This suggested that the initial biomass of *U. prolifera* was approximately 4,700 tons. This approximate coverage and biomass of the initial *U. prolifera* was determined according to the survey by Liu et al. (2013) and Xu et al. (2014a). The simulation continued for 120 d from spring to the end of the summer. In this study, instantaneous environmental factors, including temperature, salinity, currents, and solar radiation intensity, were determined according to corresponding positions where the particles were floated from the physical ECS FVCOM model.

Seven particle-tracking simulations were conducted. One hundred particles were released at a location that matched the drifter's  $in_-situ$  deployment position, and the horizontal random diffusion coefficient  $(K_r)$  was set as 50 m<sup>2</sup>/s. In addition, the depth for surface and subsurface drifters were set as 0.5 m and 2 m, respectively. Thus, for these drogued drifters, only half of the buoy was exposed above the sea surface only half of the buoy was exposed above the sea surface for these drogue-drifters, and the direct wind factor was not considered in the tacking simulations.

For the wind-exposed drifters floated at the sea surface, the wind drift is one of dominating contributions of the transportation. Dagestad and Röhrs (2019) conducted drifting buoy experiments and found that windage accounted for 3% of Stokes drift. Whiting et al. (2020) chose a constant 3% coefficient in the free-floating macroalgal trajectory simulation. The setting of this coefficient was based on the debris drift simulation of the 2011 Japan tsunami (Maximenko et al., 2018). Additionally, Jones et al. (2016) reported that the horizontal movement of surface oil slicks is drifted by ~3.5% wind speed, including a 2% direct wind drag and a 1.5% wind speed adding to the surface Stokes drift (Abascal et al., 2009). The movements of free-floating macroalgae are influenced by wind and windage, which depend on the physical characteristics of drifters. The hydrodynamic surface layer had already accounted for wind movement, and the other wind drag for particle drift was composed by direct windage.

Based on the previous studies described above, the *U. prolifera*-induced total drifting windage was in a range of 2.7–3.5%. A series of particle tracking experiments were conducted in this windage range, with an interval of 0.1%. Meanwhile, one

experiment without the direct wind factor was also undertaken for reference. Totally ten groups, with 1192 particles in every group, separated with a 0.02° horizontal resolution, were deployed in batches in the particle release zone (Fig. 3b) on May 1, 2014. These particles were traced for 120 days.

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Most importantly, two realistic dynamic growth simulations were conducted. To verify the general applicability of the model, we simulated the growth and drift processes of *U. prolifera* in the YS in 2014 and 2015, respectively, with identical model configurations. In the two simulations, each particle represented 10--tone biomass of floating *U. prolifera*, so that 4,800 tons were deployed initially. The initial coverage and biomass of the *U. prolifera* were determined based on the field surveys by Liu et al. (2013) and Xu et al. (2014a). The simulation time was 135 days from April 16 to August 29. The initial particles deployed continuously from April 16 to May 15. Daily 160-ton biomass was spatial randomly released in the hot-spot zone (Fig. 3b) over an entire month. The horizontal random diffusion  $K_r$  of the green tide simulation was set as 200 m<sup>2</sup>/s in the green tide simulation. In this study, instantaneous environmental factors, including temperature, nutrients, solar radiation intensity, ocean flow, and wind speed, were determined from the physical ECS-FVCOM model.

We set the parameters of the ecological module according to the physiological characteristics of *U. prolifera*. The functions of temperature were determined referring to the results of laboratory studies. U. prolifera has the optimal photosynthesis 325 efficiency at 20 °C, and turns white and declines rapidly under high temperature and high light intensity (Cui et al., 2015; Song et al., 2015). When the temperature is suitable (5–25.7 °C), the temperature limitation of photosynthesis  $f_n(T)$  and respiration  $f_r(T)$  are consistent. When the temperature becomes unsuitable (<5 °C or >25.7 °C), the respiration of macroalgae, unlike photosynthesis, will remain at a lower level. When the high temperature exceeds a suitable situation (>25.7 °C), the mortality 330 process replaced photosynthesis as dominance.

$$f_p(T) = \begin{cases} -4.942 \times 10^{-4} T^3 + 0.01885 T^2 - 0.135 T + 0.1014, & 5^{\circ} C \le T \le 25.7^{\circ} C \\ 0, & T < 5^{\circ} C \text{ or } 25.7^{\circ} C < T \end{cases}$$
(18)

$$f_p(T) = \begin{cases} -4.942 \times 10^{-4}T^3 + 0.01885T^2 - 0.135T + 0.1014, & 5^{\circ}C \le T \le 25.7^{\circ}C \\ 0, & T < 5^{\circ}C \text{ or } 25.7^{\circ}C < T \end{cases}$$

$$f_r(T) = \begin{cases} -4.942 \times 10^{-4}T^3 + 0.01885T^2 - 0.135T + 0.1014, & 5^{\circ}C \le T \le 25.7^{\circ}C \\ 0.789, & T < 5^{\circ}C \text{ or } 25.7^{\circ}C < T \end{cases}$$

$$f_m(T) = \begin{cases} 0, & T \le 25.7^{\circ}C \\ 0.01416T^3 - 1.223T^2 + 35.22T - 337.73, & T \ge 25.7^{\circ}C \end{cases}$$

$$(18)$$

$$f_m(T) = \begin{cases} 0, & T \le 25.7^{\circ}C \\ 0.01416T^3 - 1.223T^2 + 35.22T - 337.73, & T \ge 25.7^{\circ}C \end{cases}$$
 (20)

According to the floating growth characteristics of U. prolifera, the self-shading limited function  $f(\rho)$  was determined. When the assembled density does not exceed 0.16 mol C/m<sup>2</sup>, the growth of *U. prolifera* is not restricted by self-shading. However, as the density increases, the accumulation of *U. prolifera* becomes significant, and maximum when the density is greater than  $0.56 \text{ mol C/m}^2$ .

$$f(\rho) = \begin{cases} 1, & \rho \le 0.16 \\ 2.308 \exp(-2.5\rho) - 0.54705, & 0.16 < \rho \le 0.56 \\ 0, & \rho > 0.56 \end{cases}$$
 (21)

The complete list of parameters used in the ecological module of *U. prolifera* was shown in Table. 1. 335

Table. 1. Parameters used in ecological module for *U. prolifera*, modified from Sun et al. (2020).

Parameters	Description	Value	Dimension	<u>Reference</u>
$Q_{Nmin}$	Minimum N-quota	<u>25.3</u>	mmol N mol C <sup>−1</sup>	(Fujita, 1985)
$Q_{Nmax}$	Maximum N-quota	<u>108.7</u>	$mmol\ N\ mol\ C^{-1}$	(Sun et al., 2015)
$Q_{Pmin}$	Minimum P-quota	0.097	$mmol\ N\ mol\ C^{-1}$	(Sfriso et al., 1990)
$Q_{Pmax}$	Maximum P-quota	<u>1.4</u>	$mmol\ N\ mol\ C^{-1}$	(Sun et al., 2015)
$V_{mDIN}$	Maximum uptake rate of DIN	<u>2.8</u>	$mmol\ N\ mol\ C^{-1}h^{-1}$	(Li and Zhao, 2011; Luo et al., 2012b)
$V_{mDIp}$	Maximum uptake rate of DIP	<u>0.58</u>	$mmol\ N\ mol\ C^{-1}h^{-1}$	(Luo et al., 2012b)
$K_{DIN}$	<u>Half-saturation coefficient for DIN</u>	<u>18.77</u>	$\mu mol~N~L^{-1}$	(Li and Zhao, 2011; Luo et al., 2012b)

$K_{DIP}$	Half-saturation coefficient for DIP	<u>10</u>	$\mu mol \ N \ L^{-1}$	(Luo et al., 2012b)
$P_{max}$	Maximum photosynthetic rate	<u>240.51</u>	$\mu mol~C~g~WW^{-1}h^{-1}$	(Xu et al., 2014b)
α	Photosynthetic efficiency	<u>2.52</u>	$(\mu mol\ C\ g\ WW^{-1}h^{-1})$	(Xu et al., 2014b)
			$/(\mu mol\ photons\ m^{-2}s^{-1})$	
$R_d$	Dark respiration rate	<u>18.4</u>	$\mu mol~C~g~WW^{-1}h^{-1}$	(Xu et al., 2014b)

#### 3 Results

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#### 3.1 Variations of environmental factors

#### 3.1.1 Surface wind

The wind vectors at 10-m-height near Subei coast and Qingdao coast, retrieved from ECMWF, were showed in Figure 4. -From May to July 2014, southerly and southeasterly winds prevailed in the coast of Subei and Qingdao, and the mean wind speed reached 5 m/s. However, the southerly wind was stronger throughout May in Spring. In June, southeast winds blew in the Subei coast significantly, and the Qingdao coast was still dominated by southerly windsoutherly wind still dominated the Qingdao coast, the wind speed was slightly lower than that in May. In August, the northeast wind was strengthened, especially from August 1–10.

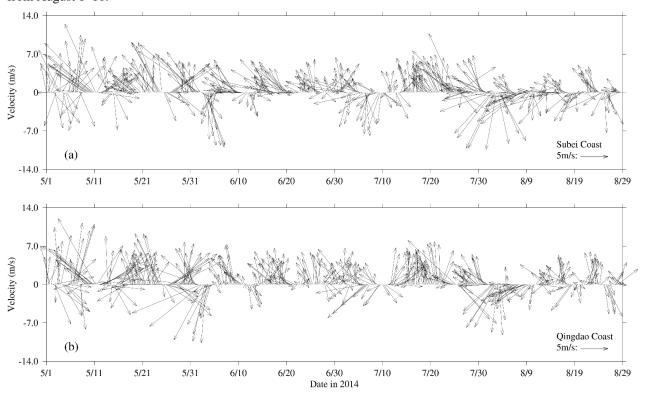


Figure. 4. Wind vectors at 10 m height near Subei coast (a) and Qingdao coast (b) at a time interval of 6 h from May to August 2014.

# 350 3.1.2 Ocean circulation

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The time-averaged distributions of surface ocean circulation every 15-d period in 2014 and 2015 are shown in Figs. 5 and 6, respectively. Affected by the southerly and southeasterly winds (Fig. 4), the coastal surface seawater flowed northward and was transported to the east of South YS. This phenomenon was more pronounced in May and late June, and July (Fig. 5a, b, d, f), indicating the possibility of *U. prolifera* drifting from the Subei Shoal toward the north. The same phenomenon was observed from May–June; and early August of 2015 (Fig. 6a–d, g). Most of the time, surface seawater from north YS was transported to South YS through the east of Rongcheng (RC). In early June and July, and August 2014 (Fig. 5c, e, g, h), the

surface seawater circulated counterclockwise in the middle region of South YS. Simultaneously, the weak currents on the south side of the Shandong Peninsula may have caused *U. prolifera* to stay in this region and gradually land on the shore. Similar ocean circulations appeared in July and late August of 2015 (Fig. 6e, f, h), and weak northward currents were observed in theon Subei Shoal.

# 3.1.3 Temperature

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Figures 5 and 6 also included the distribution of surface seawater temperature in 2014 and 2015, respectively, with every 15-d time averaging. The average temperature of South YS reached 13 °C in May (Fig. 5a, b) and increased continuously in June (Fig. 5c, d). Surface seawater temperature along Jiangsu Coast and the East China Sea was generally 1–2 °C higher than that in other areas of South YS. However, most of South YS reached a high\_temperature-state, with over 25 °C, by July 2014 (Fig. 5e, f). From mid-July to end-August, the surface temperature in Jiangsu Coast and parts of Shandong Peninsula Coast remained above 27 °C (Fig. 5f–h). The offshore sites of Qingdao and Subei were selected to determine the time series process for the physical factors (Fig. 5i–j and Fig. 6i–j). The surface temperatures of the two stations, the northern Jiangsu Coast and Qingdao coast, were increased until they reached their peaks at the end of July with over 27 °C and remained until the end of August. The distribution and tendency of South YS seawater temperature in 2015 (Fig. 6) were similar to those in 2014 (Fig. 6). However, compared with those in 2014, they had more extensive high-temperature high\_temperature coverage for South YS in August 2015 (Fig. 6g–h). The surface temperature of most YS regions exceeded 27 °C, part of the Jiangsu Coast even reached 29 °C. In addition, the surface temperatures of the two stations reached 25 °C in 2015, approximately one week later than they did in 2014 (Fig. 6i–j).

#### 375 3.1.4 Irradiation and salinity

Solar irradiation intensity is significantly different in the day and night. Therefore, only the irradiation intensity at  $\underline{\text{noon}}$  12:00 am was analyzed in Fig. 5i–j and Fig. 6i–j. Affected by the thickness of clouds, irradiation intensity at noon fluctuated drastically within 3, 200  $\mu$ mol·m<sup>-1</sup>·s<sup>-1</sup>. Compared with May and June, the irradiation intensity in July and August of 2014 and 2015 decreased slightly.

Because the Qingdao is far from the Changjiang Estuary (CJE) where the freshwater discharge enters the shelf region, the seawater salinity at Qingdao exhibited weak variations, which were maintained within the range of 31.5–32.0 PSU (Fig. 5i and Fig. 6i). However, the seawater salinity around Subei coastal region was significantly affected by the influence of the fresh water that flowed into Subei sea in summer. It dropped from approximately 32.0 PSU to approximately 30.5 PSU from May to the end of August 2014 (Fig. 5j) and to 30.0 PSU in July and August of 2015 (Fig. 6j). The surface salinity of South YS fluctuated between 29 PSU and 33 PSU during the period of the green tide bloom (Fig. 5i–j and Fig. 6i–j), which was suitable for *U. prolifera* growth (Xiao et al., 2016). For this reason, the salinity limitation was ignored in the biological module.

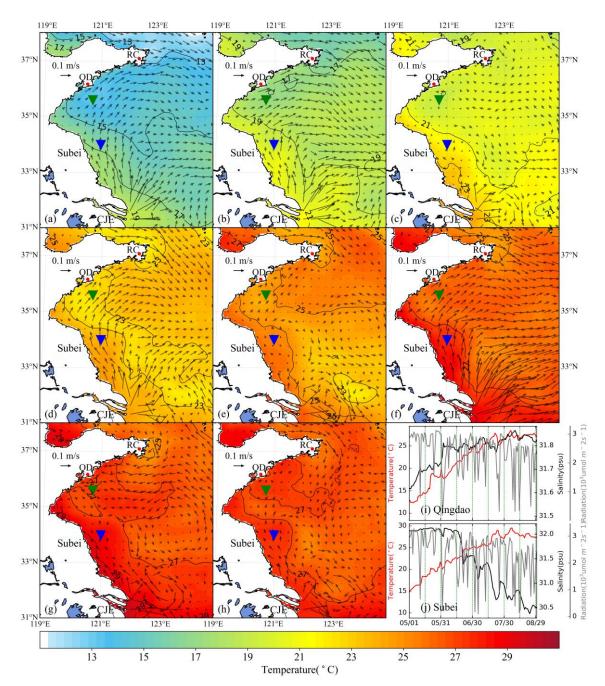


Figure. 5. Time-averaged distribution of surface current and temperature of every 15-d duration in 2014: (a) 1–15 May, (b) 16–30 May, (c) 31 May–14 June, (d) 15–29 June, (e) 30 June–14 July, (f) 15–29 July, (g) 30 July–13 August, (h) 14–29 August.

The green and blue inverted triangle indicate the position of selected Qingdao (QD) coast and Subei offshore sites, respectively. Time series of surface temperature, salinity, and irradiation in Qingdao (i) and Subei (i).

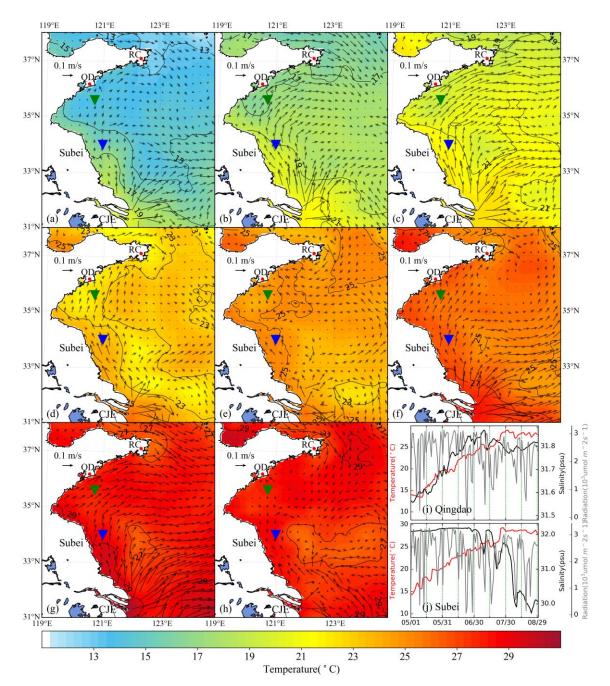


Figure. 6. Same as in Fig. 5. Time-averaged distribution of surface current and temperature of every 15-d duration in 2015: (a) 1 15 May, (b) 16 30 May, (c) 31 May 14 June, (d) 15 29 June, (e) 30 June 14 July, (f) 15 29 July, (g) 30 July 13 August, (h) 14 29 August. The green and blue inverted triangles indicate the position of selected Qingdao (QD) coast and Subei offshore sites, respectively. Time series of surface temperature, salinity, and irradiation in Qingdao (i) and Subei (i).

#### 3.1.5 Dissolved nutrients

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The dissolved inorganic nutrients in the offshore region are mainly influenced by terrestrial sources, with prominent seasonal characteristics. The concentration of dissolved inorganic nutrients in the Jiangsu region was significantly higher than in other areas. The nitrate concentration in the offshore region of Jiangsu was generally above 2 mmol/m³ in spring (Fig. 7a) and summer (Fig. 7b), especially in the Yancheng region; the nitrate concentration was over 8 mmol/m³ in spring. The nitrate concentration in the other areas of YS, except the offshore region of Jiangsu, was mainly below 2 mmol/m³. The phosphate concentration in the offshore region of Nantong and Sheyang River Estuary was still high, more than 0.6 mmol/m³ in spring (Fig. 7c) and summer (Fig. 7d). In the north of the Yancheng offshore region, phosphate concentration decreased to ~0.2

mmol/m<sup>3</sup> in summer (Fig. 7d). In the central YS and south offshore area of Shandong Peninsula, phosphate concentration was higher in summer, over 0.2 mmol/m<sup>3</sup>, than ~0.1 mmol/m<sup>3</sup> in spring.

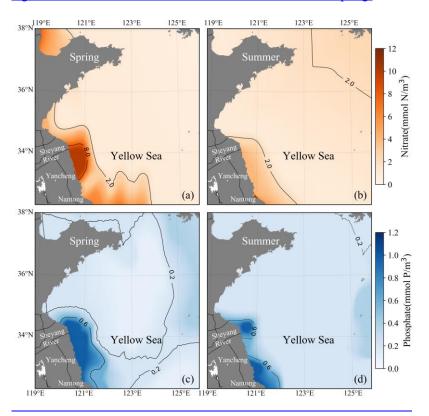


Figure. 7. Seasonal-averaged surface distributions of nitrate (upper) and phosphate (lower) in YS during spring (left), summer (right).

# 3.2 Validation of Tracking module

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## 3.2.1 Tracking module evolution

The simulated particle trajectories are generally consistent with the observed drifter trajectories, particularly for a short-term prediction (Fig. 8). The tracking time for surface drifters #1–5 lasted more than one month (Fig. 8a–e). The results show that the model was robust to reproduce the overall drifter's movement directions. Since the drifter paths may change with strong randomness due to complex variations of ocean flow, winds, and waves, the long-term prediction of drifter paths and dispersal could be of great challenge. The tracking time for subsurface drifters #6–7 (Fig. 8f) was relatively shorter, only 5–9 days. Compared with surface drifters, subsurface drifters were driven by a more complicated forcing relating to the extensivelarge depth range of the sail. The water flow at a 2-m depth was selected as the driving force approximately, considering the average drifting state of subsurface drifters. The simulated trajectories were similar to the observed drifter's movement trends. However, the drifted distance was slightly shorter than the actual situation.

The model-data compression suggests that the particle tracking algorithm algorithm in FMGDM can provide reasonable predictions for free-floating drifters with higher confidence for surface drifters. In addition, the hydrodynamic model, ECS-FVCOM, is reliable. Our results showed that *U. prolifera* were mainly under a free-floating state at the sea surface.

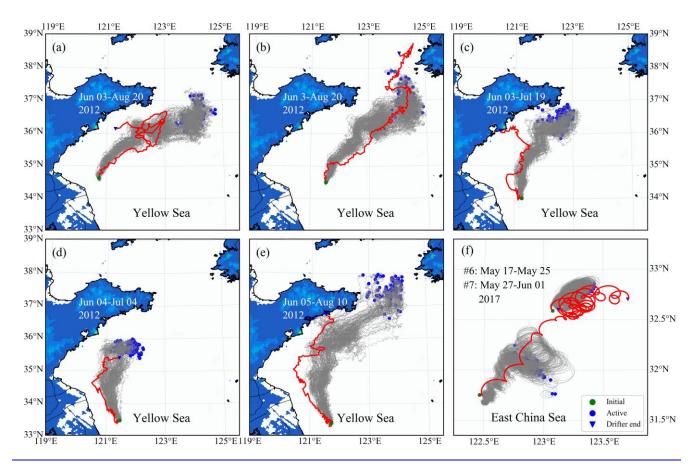


Figure. 8. Comparisons between observed (red lines) and simulated (grae-y lines) drifter trajectories. Panels (a–e) correspond to the drifters #1–5, respectively. Panel (f) corresponds to the drifters #6–7.

# 3.2.2 Windage simulation

The simulation results with different windages showed that particles first flowed northward and then turned northeastwardnorth-eastward (Fig. 9). With greater windage, the trend of northward transportation is more prominent. The particles with a windage of about 3.4% could reach the southern coast of the Shandong Peninsula on June 15 (Fig. 9b) and then turned northeastwardnorth-eastward near 124°30.00'E (Fig. 9c–d). The particle group was split at the end of July. One part drifted northward continuously to 38°N and reached the North Korean coast, and the other was turned west (Fig. 9f). The particles with a windage coefficient less than 3.2% stranded near the southern coast of the Shandong Peninsula in July and August (Fig. 9d–f). The particles without direct windage have significantly slow drifting, northernmost nearly to the south coast of Shandong Peninsula. Some particles moved northeastward north-eastward to the centralcenter of the South YS. From the comparison, winds contributed significantly to the transport of free-floating drifters. The transport results with a 2.7–3.5% windage range did not show a significant difference in the short-term simulation of 1–1.5 months (Fig. 9a–c). However, as the simulation time lasted longer, the transport pattern showed a noticeable difference (Fig. 9d–f). Compared with the evolution of green tides in the YS from remote sensing (Hu et al., 2019), it can be confirmed that the windage in a range of 3–3.2% could be applied to the drift of green tide. In this study, 3.2% was selected as the windage κ of the YS green tide simulation.

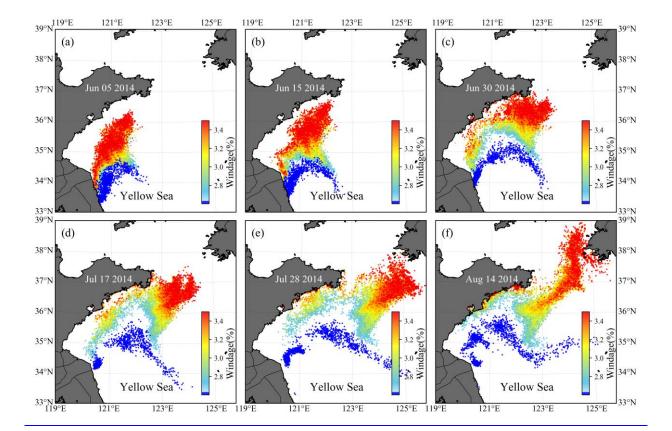


Figure. 9. Evolution of the particle distribution in FMGDM tracking simulations. Colors indicate different windage (2.7%-3.5%) of the simulated particles. The blue particles indicate the simulation without direct windage. (a) June 5, 2014; (b) June 15, 2014; (c) June 30, 2014; (d) July 17, 2014; (e) July 28, 2014; (f) August 14, 2014.

# **3.2 Effects of physical factors**

# 3.2.1 Wind

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The drift trajectory of the floating *U. prolifera* was simulated from May 1 to August 29. Three groups of initial particles were released in three regions of the Subei Shoal (Fig. 3b). In this experiment, the roles of both wind and ocean circulation were considered. Results of the experiment showed that the particles significantly drifted northward in May and then turned northeastward under the influence of south and southeast wind (Fig. 7a-c). The particles released along the coast of Subei could drift to the south side of the Shandong Peninsula in mid June and float on the north of YS throughout August (Fig. 7a). Similar to Fig. 5a, the particles released from the north side of Subei Shoal were transported northeast, and most of the particles cannot drift out of north YS (Fig. 7b). The particles released from the south side of Subei Shoal demonstrated a stronger tendency to drift eastward, which could drift to the west coast of the Korean Peninsula (Fig. 7c). These drift patterns were all a result of the full effects of wind and circulation.

### 3.2.2 Ocean circulation

The other experiment was conducted with the exclusion of the wind effect, meaning that the particles are driven only by the ocean flows. Initially, six particles were released into Subei Shoal on May 1, 2014 (lime green stars in Fig. 3b). After one and a half months of simulation, simulated particles were still floating within Jiangsu Coast and then turned northeastward under the role of the circulation to the central part of YS in July and August (Fig. 7d). Simulation results suggested that floating green tides cannot be transported out of Jiangsu Coast within a short period, but these particles still tended to move northward via wind driven current. The drift trajectory near the initial position of particle release showed obvious spiral oscillations caused by rotary tidal current. When the particles moved to the shore or north Subei Shoal, the trajectories showed small scale

465 north-south oscillation that may be attributable to alternating tidal current. However, the net transport made by periodical alternating tide current was quite limited, resulting in slow northward movement of surface particles. These two experiments with/without wind suggest the offshore extension of *U. prolifera* from the Subei Shoal to the central YS was primarily caused by the wind-induced drift. Additionally, the ocean circulation primarily caused the northward drifting of *U. prolifera*. Therefore, under the jointed effects of circulation and wind, the *U. prolifera* could detach quickly from *Porphyra* aquaculture raft from Subei Shoal.

Figure 7. Drift trajectories from wind circulation (a c) and circulation only (d) experiments. Green bullets indicate the initial positions of particles. Drift trajectory simulation for particles released along coast (a), on northern (b), and southern (c) of the Subei Shoal. Different colors are used to denote 15 d sub-periods from May 1 to August 29, 2014.

## 3.3 Simulation of dynamic growth model

# 3.3.1 Biomass of green tide

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As there was no direct way of quantifying the floating *U. prolifera* biomass of green tides throughout the YS (Wang et al., 2018), the estimated biomass data of *U. prolifera* retrieved from remote sensing observations (Hu et al., 2019) was adopted to validate the simulated biomass (Fig. 8). Satellite observations of a small amount of floating *U. prolifera* in YS were made in mid May. Until mid June, the estimated biomass of *U. prolifera* rose rapidly and peaked on June 18, 2014 (Fig. 8a), and June 21, 2015 (Fig. 8b), separately, with maximum values of 0.92 million tons and 1.98 million tons, respectively. After peaking, the biomass declined rapidly, and *U. prolifera* almost died off at the end of July. The observed biomass trends in 2014 and 2015 were similar, showing an inverted V shape.

After the initial particles were released, the green tide simulation began, with the coupling between physical drifting and biological growth. Compared with observation results, the biomass of simulation peaked after 12 d and 10 d, respectively, with a higher maximum value, 1.77 million tons on July 30, 2014 (Fig. 8a), and 2.31 million tons on June 1, 2015 (Fig. 8b). The growth trends were similar to the observations. Considering the highly random distributions, as well as the robust dynamic life history of *U. prolifera*, our simulation provided reasonable modeling results of biomass.

490 Figure 8. Comparison between simulated (red) and observed (green) biomass of green tide in 2014 (a) and 2015 (b). observed values were sourced from the study by Hu et al. (2019).

#### 3.3.21 Spatiotemporal variation of *U. prolifera*

After being released into the Subei Shoal, the initial particles drifted and dispersed by ocean flows and wind. The simulation result of the green tide in 2014 is shown in Fig. 10. It showed a small amount of *U. prolifera* floating on the Subei coast in mid-May (Fig. 10a). However, it was difficult to be observed using remote sensing technology in the early stage of green tide bloom. After one month of simulation, the modeling biomass increased to approximately 0.2 million tons (Fig. 10i). Both the result of observation and simulation showed that *U. prolifera* was transported northward and floated between northern Jiangsu offshore and Shandong Peninsula (Fig. 10b). On June 15 (Fig. 10c), both the results of observation and simulation shows that green tides had landed on the southern coast of Shandong Peninsula, including Rizhao (RZ) and Qingdao shore. Moreover, the observations shows that the green tides bloomed around the coast of Nantong (NT) and Yancheng (YC), suggesting the continuous supply of additional *U. prolifera* from aquaculture raft between May and June 2014. On June 23 (Fig. 10d), the result of both observation and simulation were consistent and showed that green tides had landed on the Shandong Peninsula on a large scale, and the farthest *U. prolifera* reached the Rushan (RS) coast. The entire coast and offshore regions were covered with a massive huge amount of floating *U. prolifera*. Due to the high concentration of nutrients, there were a large number of simulated particles growing at the Sheyang River estuary region in the entire June. The biomass of simulation

reached a peak of 0.85 million tons on June 30, and the number of simulation particles reached approximately 71,000 (Fig. 10i). Subsequently, *U. prolifera* died out quickly, and its coverage decreased significantly. On July 17 (Fig. 10f), the floating *U. prolifera* still gathered on the south coast of the Shandong Peninsula. Different from the observation result, there were some small patches of simulation result drifted eastward and reached 123°E.

In contrast with the simulation results, observation showed the re-occurrence of a large-scale green tide in Yancheng and Nantong regions from July 17 to July 28 (Fig. 10f–g), which, however, was uncaptured by the model. Both observation and simulation results showed that floating *U. prolifera* drifted eastward but still covered the south coast of the Shandong Peninsula at the end of July (Fig. 10g). After half a month, floating *U. prolifera* had died out (Fig. 10h). Observation shows that only the southern coast of Qingdao and the Subei Shoal had a few patches of *U. prolifera* on August 14, which suggests there is still a possible *U. prolifera* source near the Subei Shoal even in the summertime of July and August. However, the *U. prolifera* of simulation had almost vanished because the *U. prolifera* source was only initialized from mid-April to mid-May. The floating patch on Qingdao coast could be simulated, and some *U. prolifera* on the southern coast of Rongcheng survived. At the end of August, green tides in YS almost vanished, which could not be detected by the satellite remote sensing (Fig. 9h). The biomass of the simulation was less than 10 thousand tons.

As there was no direct way of quantifying the floating *U. prolifera* biomass of green tides throughout the YS (Wang et al., 2018), the estimated biomass data of *U. prolifera* retrieved from remote sensing observations (Hu et al., 2019) was adopted to validate the simulated biomass (Fig. 10i). The estimated biomass of *U. prolifera* rose rapidly and peaked with maximum values of 0.92 million tons on June 18, 2014 (Hu et al., 2019). The biomass declined rapidly after reaching its peak, and *U. prolifera* almost died off at the end of July. Compared with observation results, the biomass of simulation peaked after 12 days with a similar value. The growth trends between observation and simulation were similar. Considering the highly random dispersion, as well as the dynamic life history of *U. prolifera*, our simulation provides reasonable modeling results of biomass and spatial coverage.

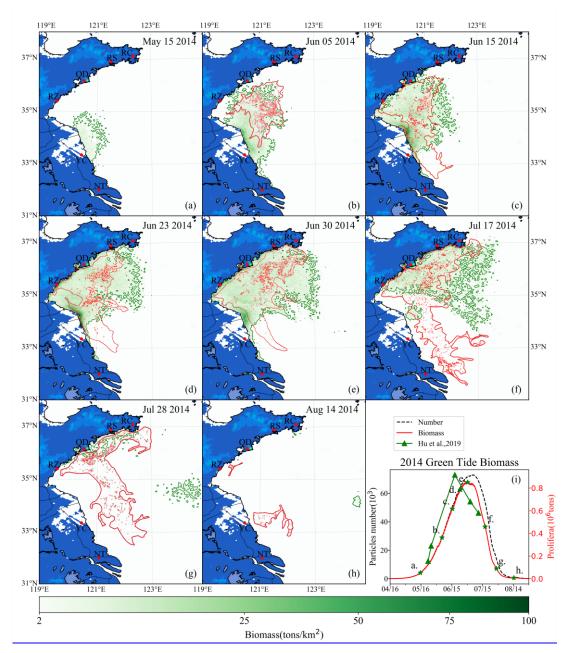


Figure 910. Comparison between simulation and remote sensing observation of green tides from May to August 2014 (a–h). The green image indicates the simulated biomass density of the green tides tide (color image, unit: tons/km²). , and tThe red image indicates shows the satellite-derived spatial coverage of *U. prolifera* from MODIS-TERRA. Panel (i) is the time series of simulated biomass and particle number of green tides in 2014, compared with observed biomass from Hu et al. (2019).and Tthe green pentagram in panel (i) indicates the biomass of the corresponding date in the panel (a–h).

To verify the reliability of the coupled model system, the green tides that bloomed in 2015 were also simulated and compared with the observations made. The simulation shows that a small amount of *U. prolifera* was floated near the coast of Jiangsu in mid-May (Fig. 11a). On May 30, the coverage of floating *U. prolifera* increased, while the northernmost green tide patches was were closed to 35°N (Fig. 11b). On June 23 (Fig. 11c), both the results of observation and simulation showed that the green tides had entered the Shandong Peninsula with large-scale coverage, distributed in most of the seas from Subei to the Shandong Peninsula and bloomed strongly offshore of Qingdao to RS. On July 2, both observation and simulation showed that the green tide still gathered along the south coast of the Shandong Peninsula, and the northernmost of the distribution range reached RC (Fig. 11d). In addition, observation showed scattered patches of *U. prolifera* floating in the center of the South YS<sub>7</sub> from June 23 to July 2, which cannot be simulated. On July 16 (Fig. 11e), satellite observations showed that the coverage

of green tide reduced greatlyconsiderably, and the distribution range was shrunk toward the west of 121°30°E. In late July, the result of simulation result showed that the biomass declined rapidly—; however, the green tides are still widely distributed in the southern regions of the Shandong Peninsula. On August 5, observations showed small patches of the floating green tide in the middle of South YS (Fig. 11g). In simulation results, a small amount of the green tide remained along the coast of the Shandong Peninsula. On August 20 (Fig. 10h11h), the green tide in the of YS completely disappeared from satellite observation and numerical simulation. Compared the biomass of observation and simulation, the estimated biomass of U. prolifera based on remote sensing peaked with maximum values of approximate 1.77 million tons on June 21, 2014 (Hu et al., 2019), while the simulated biomass peaked after 13 days with a similar value of 1.6 million. The, and the number of simulation particles peaked at with approximately 134,000 (Fig. 11i).

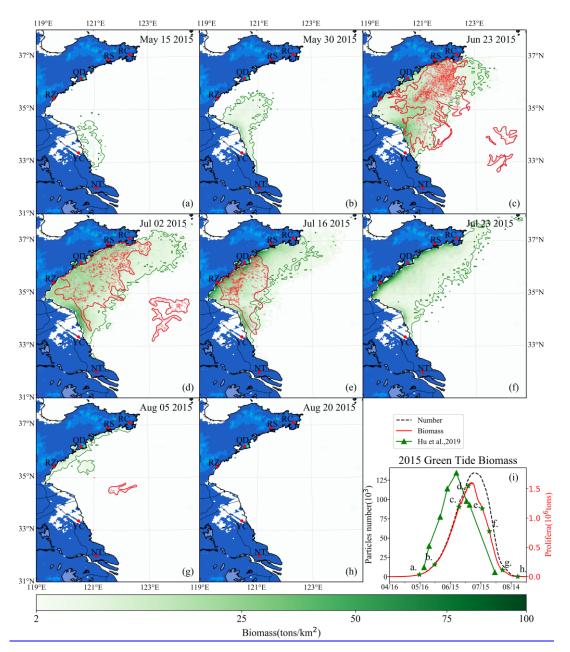


Figure 1011. Same as in Fig. 10. Comparison between simulation and remote sensing observation of green tide from May to August 2015 (a h). The green image indicates the simulated biomass density of green tide (color image, unit: tons/km²); the red image indicates the satellite derived spatial coverage of *U. prolifera* from MODIS TERRA. Panel (i) is a time series of simulated biomass of green tide in 2014, the green pentagram indicates the biomass of the corresponding date in panel (a h).

#### 4 Discussion

#### 4.1 Uncertainties of physical, biological, and anthropic processes

The observation of the entire bloom process is technically complex in the study of massive floating macroalgal blooms. <u>In this study</u>, a floating macroalgae growth and drift model was established, supplemented by remote sensing observations, which can reproduce the entire process and predict the development of macroalgae bloom. Following the establishment of floating macroalgae growth and drift model in this study, and supplemented by remote sensing observations, the entire process of growth and drift under floating could be re-produced and bloom development predicted. However, there are many uncertainties throughout the blooming process from early spring to late summer, which could significantly limit the precision of long-term prediction.

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In the realistic numerical simulations of green tides from 2014 to 2015, the initial biomass of *U. prolifera* had the same configuration of 4,800 tons,—deployed continuously from April 16 to May 15 base on the estimation in previous studies (Liu et al., 2013; Xu et al., 2014a). The initial distribution was also uniform. However, high uncertainties regarding the biomass and distribution were observed. The initial biomass of *U. prolifera* was determined primarily by the scale of local *Porphyra* aquaculture around the Subei coastal region and the timing of harvest activities. The precise estimation of initial biomass and timing requires extensive monitoring for these activities, as well as robust and timely satellite assessment of satellite remote sensing.

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From the satellite observations in June and July 2014, we observed stable patches of *U. prolifera* off the Subei Shoal (Fig. 10c–g). This indicatinged the continuous supply of *U. prolifera* from the local *Porphyra* aquaculture activities in summer, resulting in stable bloom off the Subei Shoal and northward drift. Therefore, this factor, which could lead to <u>significantmajor</u> bias of *U. prolifera* distribution and biomass, should be considered during long-term simulation.

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During the bloom of green tide blooms, large-scale salvage operations were implemented to reduce the biomass of floating U. prolifera in Jiangsu and Shandong coastal waters (Liu et al., 2013; Wang et al., 2018), which. This could significantly change the local biomass. The biomass of salvage operations reaches  $1.5-2.0 \times 10^6$  tons every spring and summer along the Shandong coastal region (Ye et al., 2011; Zhou et al., 2015), which could be the reason for the underestimation of biomass from June 2–16, 2015 (Fig. 10d–e). The salvage operations cause significant uncertainty for numerical prediction, particularly along the coast where the operations are primarily conducted.

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The propagules are distributed near the floating *Ulva* with a high density and move together with ocean flows (Li et al., 2017). The modified clay (MC) at a proper dose can flocculate with microscopic propagules and effectively remove microscopic propagules from the water column (Li et al., 2020). The physiological processes of *Ulva* cells could be disrupted by MC (Zhu et al., 2018). This method was frequently used to mitigate blooms in local areas (Li et al., 2017). The intervention of human activities on the blooming process was not considered in the model. <u>Large-scale salvage and elimination activities play essential roles in reducing the scale and intensity of the green tide bloom.</u> When the observed biomass peaked, the biomass in the simulation maintained an increasing trend. <u>Finally, The the maximum simulated biomass was larger more significant thansimilar with</u> the maximum estimated biomass<sub>2</sub>, and the bloom duration duration of the bloom was longer than that of the actual condition. <u>Large scale salvage and elimination activities play essential important roles in reducing the scale and intensity of the green tide bloom.</u>

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## 4.2 Short-term variations and quick response

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To reduce the errors of long-term simulation, caused by the complex origin of initial floating macroalgae and the uncertainty of growth and drift, the time of each simulation was shortened by dividing the entire long-term simulation into multiple short-term simulations and renewed the location and biomass in every short-term modeling by initialization of floating estimated by remote sensing observation.

Two consecutive simulations were carried out during the heyday of the YS green tides. One was configured for simulations from June 15–23, 2014, and the other from June 23–30, 2014. According to the distribution from remote sensing observation and estimated biomass from Hu et al. (2019), the initial biomass and distribution of *U. prolifera* on June 15, 2014, was determined as shown in Fig. 12a and June 23, 2014, in Fig. 12d. The initial biomass was approximately 0.7–83 million on June 15 and approximately 0.96 million on June 23.

The time interval between two consecutive <u>cloud-freeelear</u> satellite observations of green tides was generally large.\_-Two intermediate results between the <u>gap of</u> satellite observation <u>gap</u> were shown in Figs. 12b and 12e. After nearly one\_week of simulation, the coupled model system made precise simulation, compared with remote sensing (Fig. 12c, f), and the biomass was similar to that estimated via satellite remote sensing (Fig. 12g). Moreover, spatial distribution was well predicted. Compared with long-term simulation, the variation of green tide distribution and biomass could be determined more accurately by the results of the short-term simulation. The accuracy of short-term simulations is reliable, and the short-term prediction of floating macroalgal blooms can be achieved by combining the numerical model with the the satellite observations.

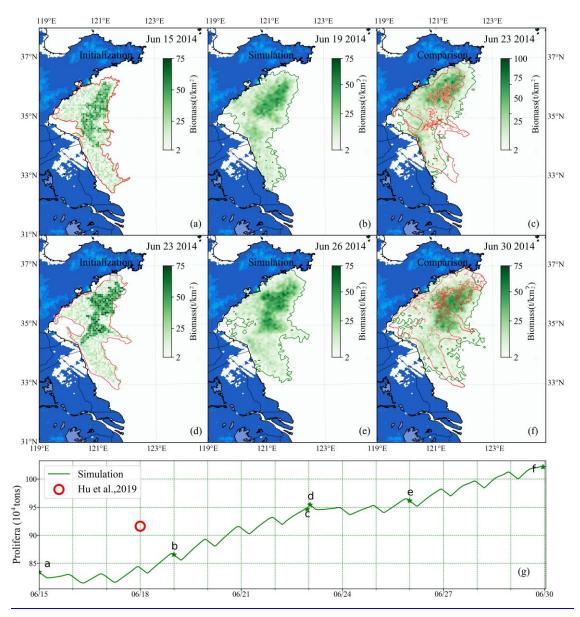


Figure 12. Comparison of short-term simulation and satellite observation of green tides in 2014: position and distribution of green tide density released initially: 15 June (a) and 23 June (d) based on the satellite date data and estimated biomass. Panels (b) and (e) are the modeling results of 19 June and 26 June, respectively. Panels (c) and (f) are the comparison between compare simulation and satellite observation on June 23 and June 30, respectively. The green image indicates the simulated result, and the red one indicates satellite observation. Panel (g) is the comparison of two consecutive simulated biomass (green line) and estimation biomass of satellite data (red circle). The green pentagrams indicate the biomass of the corresponding date in panels (a–f).

## 4.3 Roles of initial biomass and nutrient limitation

The existence of diverse origins and continuous input of floating propagules greatly significantly challenge the precise prediction and effective control of massive floating macroalgal blooms. In addition to the extensive large provision from *Porphyra* aquaculture rafts in the Subei Shoal, the somatic cells, indicated by a laboratory study, could overwinter and restore growth on the annual spring bloom (Zhang et al., 2009), which is another significant source of *U. prolifera*. Additionally, four overwintering *Ulva* propagules that existed in sediments, including *U. prolifera*, may recover their growth when the temperature and irradiation are appropriate (Liu et al., 2012). Every April, before the occurrence of green tides, *Ulva* propagules are already widespread on the southern coast of the YS (Yuanzi et al., 2014). The transport trajectory was strongly

affected by the origin of *U. prolifera*. Under the same environmental conditions, the scale of the bloom was determined primarily by the initial organisms. During the macroalgal bloom, the propagules supply from the coastal waters is continuously uncertain and difficult to determine through satellite observations or *in-\_situ* surveys. Therefore, the feature that there was still large-scale *U. prolifera* distribution around the Subei Shoal in June and July 2014 has not yet been reappeared captured, as shown by satellite observations (Fig. 910). However, as the continuous biomass entering into the ocean was not included in our simulation, this feature has not yet been determined despite the continuous entering during the period of *Porphyra* aquaculture rafts collection (mid-April to mid-May) has been considered in the green tide simulations.

Nutrient eutrophication frequently results in macroalgal blooms in coastal waters (Liu et al., 2013). The influences of nutrients in macroalgae growth have been considered in the model. Due to the difficulty in obtaining the distribution and variations of annual observational or simulated nutrients datasets, the accuracy of macroalgae simulation was limited by the deviation of nutrients datasets. In addition to the temperature, irradiation, and salinity, environmental factors that affect the growth of macroalgae, such as dissolved nutrients (Li et al., 2020a; Wang et al., 2019), were not considered in this model. Floating *U. prolifera* can efficiently absorb nutrients (Luo et al., 2012a), and the concentration of nutrients in the sea would decrease sharply when *U. prolifera* blooms dramatically, which may hinder the rapid growth of *U. prolifera* (Wang et al., 2019). When the green tide bloom reached its peak with millions of tons of biomass or drifted to the regions far from offshore, dissolved nutrient concentration may be an import significant growth limitation for growth, while. In contrast, the temperature and irradiation in the YS have not reached the limit level to large-considerable amount mortality of *U. prolifera* (Figs. 5 and 6).

## 4.4 Prospects on model development

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No technique was identified for the precise quantification of the biomass of floating macroalgae (Sun et al., 2020). Most growth models only considered the environmental factors in a fixed station and disregarded the spatial variation of floating growth. The environmental factors vary significantlygreatly at different locations. Base on Lagrangian particle tracking, each particle was considered an independent simulation unit, and td on Lagrangian particle tracking particle tracking, each particle was considered an independent simulation unit. The Each independent particle's drift velocity and growth rathe drift velocity and growth rate of each independent particle for that particle were obtained according to the natural environmental factors corresponding to the spatial position and time that particles locate. The simulation principle of this model is suitable for the actual situation of massive floating macroalgal blooms, which float and grow across vast regions.

Nutrient eutrophication frequently results in macroalgal blooms in coastal waters (Liu et al., 2013). Despite the difficulty in obtaining the distribution and variations of nutrients, Large The large-scale bloom of floating macroalgae affected the distribution of nutrients. The simulations in future studies should incorporate the circulation of nutrients between macroalgae and the ocean environment to improve the coupled model development at a more precise spatiotemporal scale. By coupling with the regional ecosystem or biogeochemical model, this model can be used to study the consumption of nutrients by the macroalgae blooms and its limitation on the growth of macroalgae. In particular, the model of floating *U. prolifera* could be established as a warning system of green tide disaster forecasting and be an efficient and economical tool for the prevention and management of green tides. Despite being used for the simulation ofto simulate the green tides, this coupled model can also be applied to other large-scale macroalgae disasters that bloom in different parts of the world.

# **5** Conclusion

A system that coupled the ecological dynamic growth module with the physical drift module for macroalgae was developed to study the spatial and temporal variations of massive floating macroalgal bloom. The dynamic process of growth and drift is

achieved by the replication/extinction and Lagrangian-based particle-tracking-of-particles. It was applied to the dynamic simulation of the YS green tide blooms that occurred-in 2014 and 2015, with environmental drivers from ECS-FVCOM. The simulation results were verified against various observation data and demonstrated reasonable prediction precision. The modeling experiments also suggested that the surface wind played a crucial role in the northward drifting of *U. prolifera* from the local-Subei Shoal, and finally resulted in an annual ecosystem disaster for the adjacent coastal region. The realistic simulation for two years exhibited many uncertainties from natural and human processes during the long duration from early spring to late summer-that could potentially lead, potentially leading to extensive prediction bias. However, the short-term simulation in this model, along with the determination of spatial coverage and biomass, proved to be an efficient and robust system for the provision of-accurate forecasting of the development of *U. prolifera*.

Although this unique tool for macroalgae prediction was only applied in the simulation of <u>the</u>-YS green tide, it can potentially be used to study other macroalgae <u>blooms</u> such as golden tides caused by *Sargassum*, in <u>other different</u> regions <u>where</u> with sufficient information on the macroalgae physiological relationship with environmental factors <u>are available</u>.

and reasonable ocean dynamics model.

# Code and data availability

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690 The Fortran code of **FMGDM** v1.0is available at https://doi.org/10.5281/zenodo.4459922https://doi.org/10.5281/zenodo.5067726 (last access: 316 July March 2021). The example of the greet green tides bloomed in the Yellow Sea, China, is available at https://doi.org/10.5281/zenodo.4607828https://doi.org/10.5281/zenodo.5067743 https://doi.org/10.5281/zenodo.4607829 (last access: 316 July March 2021). The ECS-FVCOM forcing data (surface wind, radiations), the ocean bathymetry, and the 695 results of ECS-FVCOM (water velocity, temperature, and salinity nutrients), which are also used as input variables of FMGDM for the green tides simulation, the information of seven groups initial particles position, the satellite pictures of the green tides in the YS, 2014 and 2015, and the drifter trajectories dataset used to evaluate the tracking module are available at: https://doi.org/10.5281/zenodo.4616462 http://doi.org/10.5281/zenodo.4620534 (last access: 18-7 July March 2021).

# 700 **Author contribution**

JG proposed and led this model development study. JG and FZ developed the coupled model. DL provided many important suggestions for this study, and key data of *U. prolifera* growth. JG, PD, and CC contributed to the the simulation result analysiss of the ECS-FVCOM—systems which is used s necessary for this research. XW contributed to the remote sensing results interpretation. FZ processed the model outputs and wrote the manuscript with contributions from all co-authors.

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# **Competing Interests**

The authors declare that they have no conflict of interest.

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