# A Regional multi-Air Pollutant Assimilation System (RAPAS v1.0)

2	for emission estimates: System development and application
3	Shuzhuang Feng <sup>1</sup> , Fei Jiang <sup>1,2</sup> , Zheng Wu <sup>3</sup> , Hengmao Wang <sup>1,2</sup> , Wei He <sup>1</sup> , Yang Shen <sup>1</sup> ,
4	Lingyu Zhang <sup>1</sup> , Yanhua Zheng <sup>1</sup> , Chenxi Lou <sup>1</sup> , Ziqiang Jiang <sup>4</sup> , Weimin Ju <sup>1,2</sup>
5	
6	<sup>1</sup> Jiangsu Provincial Key Laboratory of Geographic Information Science and Technology, International
7	Institute for Earth System Science, Nanjing University, Nanjing, 210023, China
8	<sup>2</sup> Jiangsu Center for Collaborative Innovation in Geographical Information Resource Development and
9	Application, Nanjing, 210023, China
10	<sup>3</sup> Chongqing Institute of Meteorological Sciences, Chongqing, 401147, China
11	<sup>4</sup> Jiangsu Environmental Monitoring Center, Nanjing, 210019, China
12	
13	
14	
15 16	Correspondence to: Fei Jiang (jiangf@nju.edu.cn)
17	
18	
19	
20	
21	
22	
23	
24	
25	
26	
27	
28	
29	

# Abstract

30

31

33

57

58

distributed observations of atmospheric compositions, which is a vital means for 32 quantifying large-scale anthropogenic and natural emissions. In this study, we developed a Regional multi-Air Pollutant Assimilation System (RAPAS v1.0) based on 34 35 the Weather Research and Forecasting/Community Multiscale Air Quality Modelling 36 System (WRF/CMAQ) model, the three-dimensional variational (3DVAR) algorithm, 37 and the ensemble square root filter (EnSRF) algorithm. It is capable of This system can simultaneously assimilating assimilate spatially distributed hourly in situin-situ 38 measurements of CO, SO<sub>2</sub>, NO<sub>2</sub>, PM<sub>2.5</sub> and PM<sub>10</sub> concentrations observations to 39 quantitatively optimizeinfer gridded emissions of CO, SO<sub>2</sub>, NO<sub>x</sub>, primary PM<sub>2.5</sub> 40 41 (PPM<sub>2.5</sub>), and coarse PM<sub>10</sub>(PMC) on a regional scale. In each data assimilation window, we use a RAPAS includes two subsystems, initial field assimilation (IA) subsystem and 42 emission inversion (EI) subsystem, which are used to generate a good chemical initial 43 44 condition (IC), and conduct inversions of anthropogenic emissions, respectively. A 45 "two-step" inversion scheme is adopted in the EI subsystem in each data assimilation 46 (DA) window, in which the emission is inferred in the first step, and then, it is input 47 into the CMAQ model to simulate the initial condition (initial fieldIC) of the next window, meanwhile, it is also. The posterior emission is transferred to the next window 48 49 as the prior emission. The chemical IC is optimized through the IA subsystem, and the original emission inventory is only used in the first DA-window. Besides Additionally, 50 51 a "super-observation" approach is implemented based on optimal estimation theory to 52 decrease the computational costs and, observation error correlations, and reduce the 53 influence of representativeness errors. 54 With Using this system, we estimated the emissions of CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and 55 PMC in December and July 2016 over China using the corresponding nationwide surface observations. The 2016 Multi-resolution Emission Inventory for China (MEIC 56 2016) was used as the prior emission. For December, the system was run from 26 November to 31 December, in which the IA subsystem was run in the first 5 days, and

Top-down atmospheric inversion infers surface-atmosphere fluxes from spatially

the EI subsystem was run in the following days. In July, the system was run in the same way. The evaluation and sensitivity testing of this system mainly focused on December. The rResults showed that compared to the prior emissions (MEIC 2016), the posterior emissions of CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC in December 2016 increased by 129%, 20%, 5%, 95%, and 1045%, respectively, and the emission uncertainties decreased by 44%, 45%, 34%, 52%, and 56%, respectively. With the inverted emissions, the simulated concentrations of CO, NO<sub>2</sub>, SO<sub>2</sub>, PM<sub>2.5</sub> and PM<sub>10</sub> with the prior inventory have large systematic biases, with relative biases in the range of -48.2-54.2%. In the IA subsystem, after 3DVAR, the root mean squared error (RMSE) of the simulated concentrations decreased by 50.0-73.2%, and the correlation coefficient (CORR) increased to 0.78-0.92 for the five species. In the EI subsystem, after emission inversions, the RMSE of the simulated concentrations decreased by 40.1\_-56.3%, and the CORR increased to 0.69-0.87. For the whole mainland China, the uncertainties were reduced by 44.4%, 45.0%, 34.3%, 51.8% and 56.1% for CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub> and PMC, respectively. Overall, compared to the prior emission (MEIC 2016), the posterior emissions increased by 129%, 20%, 5%, and 95% for CO, SO<sub>2</sub>, NO<sub>x</sub> and PPM<sub>2.5</sub>, respectively, indicating that there was significant underestimation in the MEIC inventory. The posterior PMC emissions, including anthropogenic and natural dust contributions, increased by 1045%. A series of sSensitivity tests were conducted with different inversion processes, prior emissions, prior uncertainties, and observation errors. The rResults showed that the "two-step" scheme clearly outperformed the simultaneous assimilation of ICs and emissions ("one-step" scheme) in emission inversion, and the system is rather robust in estimating the emissions using the nationwide surface observations over China. Our This study offers a useful tool for accurately quantifying multi-species anthropogenic emissions at large scales and in near\_-real time.

86 87

59

60

61

62

63

64

65

66

67

68

69

70

71

72

73

74

75

76

77

78

79

80

81

82

83

84

#### 1. Introduction

88

89

90

91

92

93

94

95

96

97

98

99

100

101

102

103

104

105

106

107

108

109

110

111

112

113

114

115

116

Due Owing to rapid economic developments and pollution control legislations, there is an increasing demand to provide updated emission estimates has arisen, especially in areas where anthropogenic emissions are intensive. Accurately estimating source emission quantities and spatiotemporal changes resulting from various regulations is imperative and valuable for understanding air quality responses and is crucial for providing timely instructions for the design of future emissions regulations. However, most inventories have beenwere developed based on a bottom-up approach and are usually updated with a delay of a few years a few years delayowing due to the complexity of gathering all-statistical information on activity levels and sector-specific emission factors (Ding et al., 2015). The large uncertainty associated with the low temporal and spatial resolutions of these datasets also greatly limits the assessment of emission changes. Some studies (Bauwens et al., 2020; Shi and Brasseur, 2020) have evaluated emission changes indirectly through concentration measurements, but; however, air pollution changes are not only dominated by emission changes, but also highly affected by meteorological conditions (Shen et al., 2021). Top-down atmospheric inversion infers surface-atmosphere fluxes from spatially distributed observations of atmospheric compositions. Recent efforts have focused on developing air pollution data assimilation (DA) systems to conduct the top-down inversions, which is able tocan integrate model and multi-source and large amounts of observational information to constrain emission sources. Two major methods, namely, 4D-variational data assimilation (4DVAR) and ensemble Kalman filter (EnKF), are widely used in those DA systems: 4D-variational data assimilation (4DVAR) and ensemble Kalman filter (EnKF). 4DVAR provides a global optimal analysis through by minimizing a cost function. It shows an implicit flow-dependent background error covariance and can reflect complex nonlinear constraint relationships (Lorenc, 2003). Additionally, a weak constraint 4DVAR methodthe model error can be partly accounted for the model error with a weak constraint 4DVAR method through the by defining definition of a systematic error term in a cost function (Derber, 1989). For example, the

GEOS-Chem and TM5 4DVAR frameworks have been used to estimate CH4 (Alexe et 117 118 al., 2015; Monteil et al., 2013; Schneising et al., 2009; Stanevich et al., 2021; Wecht et 119 al., 2014) and CO<sub>2</sub> fluxes (Basu et al., 2013; Nassar et al., 2011; Wang et al., 2019a) 120 from different satellite retrieval products. Monteil et al. (2013) showed that the global patterns of CH<sub>4</sub> emissions derived from SCIAMACHY (with bias correction) and 121 122 GOSAT retrievals are in remarkable agreement based on 15 months observations. Additionally, Jiang et al. (2017) and Stavrakou et al. (2008) Kurokawa et al. (2009) also 123 124 used the 4DVAR algorithm to\_estimate global CO and NO<sub>x</sub> emission trends from 2000 2015 using MOPITT and GOME/SCIAMACHY retrievals, respectively-125 Kurokawa et al. (2009) and Stavrakou et al. (2008) also used the 4DVAR technique to 126 estimate NO<sub>\*</sub> emission. Using NIES LiDAR observations, Yumimoto et al. (2008) 127 applied the 4DVAR DA to infer dust emissions over eastern Asia and the results agreed 128 well with various satellite data and surface observations. Based on surface observations, 129 Meirink et al. (2008) developed a 4DVAR system to optimize monthly methane 130 emissions, which showed a high degree of consistency in posterior emissions and 131 132 uncertainties when compared with an analogous inversion based on the traditional 133 synthesis approach. ICchanges. 134 Although considerable progress has been made to reduce large uncertainties in emission 135 inventories, tThe drawback of the 4DVAR method is the additional development of adjoint models, which that are technically difficult and cumbersome for complex 136 137 chemical transport models (Bocquet and Sakov, 2013). Instead, EnKF uses the flow-138 dependent background error covariance generated by ensemble simulations to map the deviations in concentrations to increments of emissions, which is more flexible and 139 140 easier to implement. Many previous studies have used EnKF techniques to assimilate 141 the single or dual species observations to optimize the corresponding emission species 142 (Chen et al., 2019; Peng et al., 2017; Schwartz et al., 2014; Sekiyama et al., 2010). 143 Miyazaki et al. (2017) improved NO<sub>x</sub> emission estimates using multi-constituent 144 satellite observations, and further estimated global surface NO<sub>x</sub> emissions from 2005 to 2014. Feng et al., (2020b) used surface observations of NO<sub>2</sub> to infer the NO<sub>x</sub> emission 145

changes in China during the COVID-19, and quantitatively evaluate the impact of the epidemic on economic activities from the perspective of emission change. Tang et al. (2011) reported that the simultaneous optimizations of the ICs of O<sub>3</sub>, NO<sub>3</sub> and volatile organic compounds (VOCs) and adjusted the emissions of NOx and VOCs through assimilating surface O<sub>3</sub> observations and achieved<del>produced</del> an overall better performance in ozoneO<sub>3</sub> forecasts than the adjustment in emissions only. However, such a revision may encounter the problem of model error compensation rather than a retrieval of physically meaningful quantities, which should be avoided from overfitting for emission inversion purposes (Bocquet, 2012; Navon, 1998; Tang et al., 2011). The EnKF has also been widely applied to optimize emissions of carbon dioxide (Jiang et al., 2021; Liu et al., 2019), carbon monoxide (Feng et al., 2020a; Mizzi et al., 2018), sulfur dioxide (Chen et al., 2019), ammonia (Kong et al., 2019), etc. Multi-species data assimilation has shown the advantage of can efficiently reducing reduce the uncertainty in emission inventories and has led to improvements in air quality forecasting (Ma et al., 2019; Miyazaki et al., 2012b), since as it would offers additional constraints on emission estimates through the improvements in related atmospheric fields, chemical reactions, and gas-particle transformations (Miyazaki and Eskes, 2013). Barbu et al. (2009) updated sulfur oxide ( $SO_x$ ) emissions with  $SO_2$  and sulfate aerosol observations and found that the simultaneous assimilation of both species had performed better performance than assimilating one of them aloneseparately. Muller and Stavrakou (2005) also found that the simultaneous optimization of the sources of CO and  $NO_x$  led to better agreement between simulations and observations compared to the case where only CO observations are used. The deviation in the chemical initial condition (IC) is one of thean important sources of error that affects the accuracy of emission inversion, because atmospheric inversion fully attributes the biases in simulated and observed concentrations to the deviations in emissions (Meirink et al., 2006; Peylin et al., 2005). The biases of concentrations would be compensated through by the unreasonable adjustment of pollution emissions without the optimization of ICs (Tang et al., 2013). Tang et al. (2011) reported that the

146

147

148

149

150

151

152

153

154

155

156

157

158

159

160

161

162

163

164

165

166

167

168

169

170

171

172

173

simultaneous optimizations of the ICs of O3, NO, and volatile organic compounds
(VOCs) and the emissions of NO, and VOCs produced an overall better performance
in ozone forecasts than the adjustment in emissions only. Similar Smethod of
simultaneously optimizing chemical ICs and emissions were also has been applied to
constraining emissions in many previous studies (Ma et al., 2019; Miyazaki et al., 2012a
Peng et al., 2018). For example, Elbern et al. (2007) adjusted O <sub>3</sub> ICs, NO <sub>x</sub> ICs and
emissions, VOCs ICs and emissions jointly through assimilating surface O <sub>3</sub> and NO <sub>x</sub>
observations. Although the forecast skills of O <sub>3</sub> were improved, due to the coarse model
resolution and the strong nonlinear relationship between O <sub>3</sub> and NO <sub>x</sub> , the assimilation
of O <sub>3</sub> observation worsened emission inversion and forecast of NO <sub>x</sub> . Peng et al. (2018)
assimilated near-surface observations to simultaneously optimize the ICs and emissions
In the 72-hr forecast evaluation, their resultant emission succeeded in improving SO <sub>2</sub>
forecast while having little influence on CO and aerosol forecast and even degrading
the forecast of NO <sub>2</sub> . Ma et al. (2019) also found that the DA benefits for forecast almost
disappeared after 72 hr using optimized ICs and emissions. Although a large
improvement has been achieved, this method still has great significant limitations in
emission inversion because as the contributions from the emissions and the chemical
ICs to the model's biases are difficult to distinguish (Jiang et al., 2017). In addition, in
this method, the constraints of the chemical ICs with observations in each assimilation
window make the emission inversions—are independent between the assimilation
windows <u>independent. This</u> , means <u>that</u> if the emission in one window is overestimated
or underestimated, it cannot be transferred to the next window for further correcting
correction and be compensationed in the following windows. Considering the
importance of emissions in chemical field prediction (Bocquet et al., 2015), the rapid
disappearance of the DA benefits seems unrealistic, indicating that simultaneously
optimizing chemical ICs and emissions This may result in a systematic bias in the
inverted emissions (Jiang et al., 2021).
Since 2013, China has deployed an air pollution monitoring network that publishes

nationwide and real-time hourly surface atmospheric observations. This dataset

provides an opportunity to improve emission estimates using the DA. In this study, a regional multi- air pollutant assimilation system introducing using 3DVAR and EnKF DA techniques is was constructed to simultaneously assimilate various surface observations (e.g., CO, SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>, PM<sub>2.5</sub>, and PM<sub>10</sub>). Considering the possible shortcomings of the simultaneous optimization method (named as-"one-step" method in this study), as metioned by Jiang et al. (2021), we adopted a "two-step" method (Sect. 3) in this system. Unlike the "one-step" method, the ICs of each DA window in the "two-step" method is were simulated using the posterior emissions of the pervious previous DA window. The capabilitiesy of RAPAS in for reanalysis field generation and emission inversion estimation was were also evaluated. The robustness of the system was also investigated with different prior inventories, uncertainty settings of the prior emissions, and observation errors. The remainder of theis paper is organized as follows: in Sect. 2, we Section 2 introduces the DA system and the observation data, and in Sect. 3, weSection 3 describes the experimental design,- Section 4 presents and discusses tThe results of the system performance and sensitivity tests are presented and discussed in Sect. 4, and followed by the conclusions in Sect. 5. Section 5 concludes the paper.

221

222

223

224

225

226

227

228

229

230

231

204

205

206

207

208

209

210

211

212

213

214

215

216

217

218

219

220

#### 2. Method and data

#### 2.1 System description

# 2.1.1 Procedure of the assimilation system

A regional air pollutant assimilation system has been preliminarily constructed and successfully applied in our previous studies to optimize the gridded CO and NO<sub>x</sub> emissions (Feng et al., 2020a; Feng et al., 2020b). Herein, the system was further extended to simultaneously assimilate multiple species (e.g., CO, SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>, PM<sub>2.5</sub>, and PM<sub>10</sub>) and officially named as—the Regional multi- Air Pollutant Assimilation System (RAPASv1.0). The RAPAS mainly includes has three components: a regional chemical transport model (CTM), which is coupled offline and used to simulate the

meteorological fields and atmospheric compositions, and the 3DVAR and ensemble square root filter (EnSRF) modules, which are used to optimize chemical ICs (Feng et al., 2018; Jiang et al., 2013b) and anthropogenic emissions (Feng et al., 2020a; Feng et al., 2020b), respectively. The introduction of 3DVAR was introduced mainly considerings its excellentgreat performance inbased on our previous study and the lower computational cost during the spin-up period in optimizing ICs. Additionally, it has been found that the 3DVAR method can obtain a better IC than the EnKF method (Schwartz et al., 2014). Based on the above three components, the RAPAS is was divided into two subsystems: namely the IC assimilation (IA) subsystem (CTM plus 3DVAR) and the emission inversion (EI) subsystem (CTM plus EnSRF). As shown in Figure 1, the IA subsystem is-was first run to optimize the chemical ICs (Kleist et al., 2009; Wu et al., 2002) for the subsequent EI subsystem. In the IA subsystem, we do not need to dDistinguish the type of source types of the model-observation mismatch error was not required for the IA subsystem. The EI subsystem runs cyclically with a "two-step" scheme. In the first step, the prior emissions  $(X^b)$  are perturbed and input into the CTM model to simulate chemical concentration ensembles. The simulated concentrations of the lowest model level are were then interpolated to the observation space according to the locations and times of the observations using the nearest-neighbor interpolation method. The pPrior emissions  $(X^b)$ , simulated observations and real observations are were entered into the EnSRF module to generate the optimized emissions  $(X^a)$ . In the second step, the optimized emissions are were re-entered into the CTM model again to generate the ICs of the next DA window. Meanwhile, the optimized emissions are-were transferred to the next window as the prior emissions. Different from Unlike the "one-step" scheme, this the "two-step" scheme needs to run the CTM model twice, which is time consuming, but canit could transfer the potential errors of the inverted emissions in one DA window to the next for further correction. The benefits of this scheme will beare further presented discussed in Section. 4.3.

232

233

234

235

236

237

238

239

240

241

242

243

244

245

246

247

248

249

250

251

252

253

254

255

256

257

258

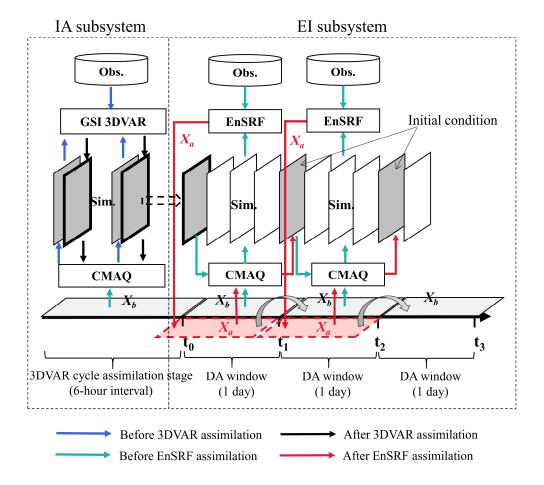


Figure 1. The eComposition and flow chart of RAPAS. The  $x_a$  and  $x_b$  represent the prior and posterior emissions. The 3DVAR assimilation stage lasts 5-five days with data input frequency of 6-six hours, and the DA window in the EI subsystem is set to 1-one day.

#### 2.1.2 Atmospheric transport model

260

261

262

263

264

265

266

267

268

269

270

271

272

273

274

The regional model of Research and chemical transport Weather Forecasting/Community Multiscale Air Quality Modelling System (WRF/CMAQ) was adopted in this study. CMAQ is a regional 3-D Eulerian atmospheric chemistry and transport model with a "one-atmosphere" design developed in by the US Environmental Protection Agency (EPA). It can simultaneously address the complex interactions among multiple pollutants/air quality issues simultaneously. The CMAQ was driven by the WRF model, which is a state\_of\_the\_art mesoscale numerical weather prediction system designed for both atmospheric research and meteorological field forecasting. In this study, WRF version 4.0 and CMAQ version 5.0.2 were adoptedused. The WRF

simulations were performed with a 36-km horizontal resolution on 169 × 129 grids, and it-covering alls the whole of mainland of China (Figure 2). This spatial resolution has been widely adopted in regional simulations and as it can provide good simulations of the spatiotemporal variations of in air pollutants (Mueller and Mallard, 2011; Sharma et al. 2016). In the vertical direction, there are were 51 sigma levels on the sigmapressure coordinates extending from the surface to 100 hPa. The underlying surface of the urban and built-up land was replaced by the MODIS land cover retrieval of 2016 to adapt to the rapid expansion of urbanization. The CMAQ model is was run with the same domain but with three grid cells removed from each side of the WRF domain. There are were 15 layers in the CMAQ vertical coordinates, which were interpolated from the 51 WRF layers. The meteorological initial and lateral boundary conditions are were both provided by the Final (FNL) Operational Global Analysis data of the National Center for Environmental Prediction (NCEP) with a 1° × 1° resolution at 6-h intervals. The chemical lateral boundary conditions and chemical ICs in the IA subsystem come originate from the background profiles. As mentioned above, in the EI subsystem, the chemical IC in the first window is provided by the IA subsystem, and in the following windows, it is forward simulated using optimized emissions from the previous window. The Carbon Bond 05 with updated toluene chemistry (CB05tucl) and the 6th generation aerosol module (AERO6) are were chosen as the gas-phase and aerosol chemical mechanisms, respectively (Appel et al., 2013; Sarwar et al., 2012). The dDetailed physical and chemical configurations are listed in Table 1.

275

276

277

278

279

280

281

282

283

284

285

286

287

288

289

290

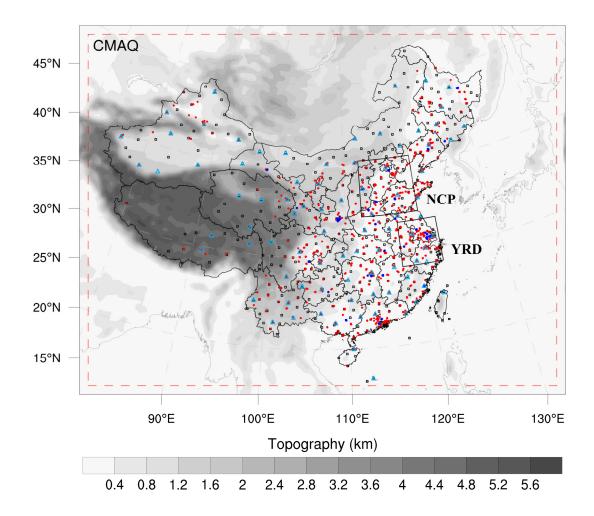
291

292

293

294

295



**Figure 2**. Model domain and observation network. The red dashed frame depicts the CMAQ computational domain; the black squares represent the surface meteorological measurement sites; the turquoise triangles represent the sounding sites; and the red and blue dots represent the air pollution measurement sites. Observations of from all sites are were assimilated in the 3DVAR subsystem, while observations of city sites where red dots are were averaged are used for assimilation and where green blue dots are were averaged are used for independent evaluation in the EI subsystem; the boxed subregions are the North China Plain (NCP) and Yangtze River Delta (YRD); and the shaded area depicts the topography.

**Table 1.** Configuration options of WRF/CMAQ

WRF		CMAQ	_
Parameter	Scheme	Parameter	Scheme
Microphysics	WSM6	Horizontal/Vertical advection	yamo/wrf
Longwave	RRTM	Horizontal/Vertical diffusion	multiscale/acm2
Shortwave	Goddard	Deposition	m3dry
Boundary layer	ACM	Chemistry solver	EBI
Cumulus	Kain-Fritsch	Photolysis	phot_inline
Land-surface	Noah	Aerosol module	AERO6
Surface layer	Revised	Cloud module	cloud_acm_ae6
Urban canopy	No	Gas-phase chemistry	CB05tucl

# 311 2.1.3 3DVAR assimilation algorithm

The Grid-point Statistical Interpolation (GSI) developed in by the US National Centers for Environmental Prediction (NCEP) was employed utilized in this study. Building upon the work of Liu et al. (2011), Jiang et al. (2013b) and Feng et al. (2018), we extended it GSI to simultaneously assimilate multiple species (including CO, SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>, PM<sub>2.5</sub>, and PM<sub>10</sub>) and first used individual aerosol species of PM<sub>2.5</sub> as analysis variables within the GSI/WRF/CMAQ framework. Additional work includes the construction of surface air pollutant observation operators, the updating of observation errors, and the statistics of background error covariance for the analysis variables. Moreover, the data interface was also modified to read/write the CMAQ output/input file directly, which is was easy to implement.

In the sense of a-minimum analysis error variance, the 3DVAR algorithm optimizes the analysis fields with observations by iterative processes to minimize the cost function (J(x)) defined below:

325 
$$J(\mathbf{x}) = \frac{1}{2} (\mathbf{x_a} - \mathbf{x_b})^T \mathbf{B}^{-1} (\mathbf{x_a} - \mathbf{x_b}) + \frac{1}{2} [H(\mathbf{x_a}) - \mathbf{y}]^T \mathbf{R}^{-1} [H(\mathbf{x_a}) - \mathbf{y}],$$
(1)

where  $\mathbf{x_a}$  is a vector of the analysis field,  $\mathbf{x_b}$  denotes is the background field,  $\mathbf{y}$  is the vector of observations,  $\mathbf{B}$  and  $\mathbf{R}$  are the background and observation error covariance matrices, respectively, representing the relative contributions to the analysis, and  $\mathbf{H}$  is the observation operator that maps the model variables to the observation space.

The analysis variables are were the 3D mass concentrations of the pollution compositions components (e.g., CO and sulfate) at each grid point. Hourly mean surface pollution observations within a one-1-hour window of the analysis are were assimilated. To assimilate the surface pollution observations, model-simulated compositions are were first diagnosed at the observation locations. For gas concentrations that are to be directly used as analysis variables, data the units need to be converted from ppm or and ppb to mg m<sup>-3</sup> and pm m<sup>-3</sup>, respectively, to match with the observations. The model-simulated PM<sub>2.5</sub> and PM<sub>10</sub> concentrations at the ground level are were diagnosed as follows:

340 
$$PM_{2.5} = f_i \times PM_i + f_j \times PM_j + f_k \times PM_k = OC + EC + SO_4^{2-} + NO_3^{-} + NH_4^{+} + PM_4^{-} + NO_3^{-} + NH_4^{-} +$$

$$341 \quad SEAS + AP_{2.5} \tag{2}$$

$$342 PM_{10} = PM_i + PM_j + PM_k = PM_{2.5} + PMC (3)$$

where  $f_i$ ,  $f_j$ , and  $f_k$  are the PM<sub>2.5</sub> fractions of the Aitken, accumulation, and coarse modes, respectively. These ratios are recommended as the concentrations of PM<sub>2.5</sub> and fine mode aerosols (i.e., Aitken plus accumulation) could can differ because the PM<sub>2.5</sub> particles include small tails from the coarse mode in the CMAQ model (Binkowski and Roselle, 2003; Jiang et al., 2006).  $PM_i$ ,  $PM_j$ , and  $PM_k$  represent are the mass concentrations of the 3-three modes in the CMAQ model, respectively. Seven aerosol species of PM<sub>2.5</sub>, including (organic carbon (OC), elemental carbon (EC), sulfate  $(SO_4^{2-})$ , nitrate  $(NO_3^{-})$ , ammonium  $(NH_4^{+})$ , sea salt (SEAS), and fine-mode unspeciated aerosols  $(AP_{2.5})$ , and additional coarse PM<sub>10</sub> (PMC) are were extracted as analysis variables, which are and were updated by using the PM<sub>2.5</sub> and PMC observations, respectively. Before the calculation of calculating equation (1) within the GSI, the

analysis variables <u>are were</u> bilinearly interpolated in the horizontal direction to the observation locations.

The computation of Calculating background error covariance (**B**) is generally costly and difficult when a high-dimensional numerical model is used. For simplification, **B** is was represented as a product of spatial correlation matrices and standard deviations (SDs).

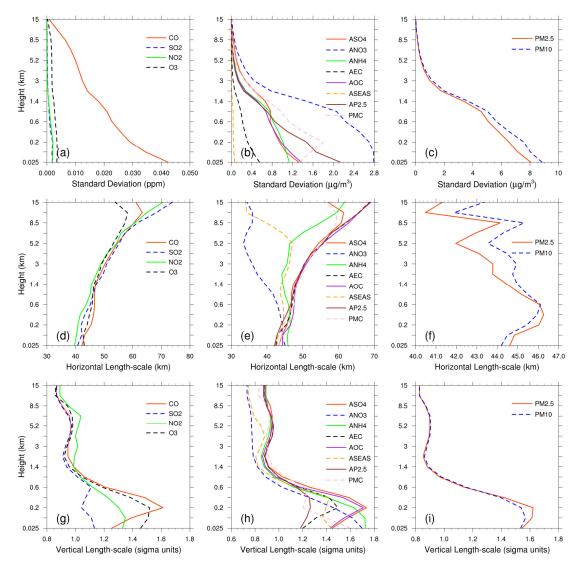
$$\mathbf{B} = \mathbf{D}\mathbf{C}\mathbf{D}^T \tag{4}$$

$$\mathbf{C} = \mathbf{C}_{\mathbf{x}} \otimes \mathbf{C}_{\mathbf{y}} \otimes \mathbf{C}_{\mathbf{z}} \tag{5}$$

where **D** is the background error SD matrix; C is the background error correlation matrix; C denotes is the Kronecker product; and  $C_x$ ,  $C_y$ , and  $C_z$  denote three one-dimensional correlation submatrices in the longitude, latitude, and vertical coordinate directions, respectively.  $C_x$  and  $C_y$  are assumed to be horizontally isotropic horizontally such that they can be represented using a Gaussian function. The correlation between any two points  $x_i$  and  $x_j$  in the horizontal direction ear be expressed as follows:

368 
$$c(x_i, x_j) = e^{-\frac{(x_i - x_j)^2}{2L^2}}$$
 (6)

where L is the horizontal correlation scale, which is estimated using the proxy of the background error (Figure 3). The vertical correlation matrixe  $C_z$  is directly estimated from the model background field since as  $C_z$  is only an  $n_z \times n_z$  (here,  $n_z=15$ ) matrix.



**Figure 3**. Vertical profiles of standard deviations (top, μg m<sup>-3</sup>), horizontal length scale (middle, km) and vertical length scale (bottom, km) length scales for CO, SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>, sulfate, nitrate, ammonium, EC, OC, sea salt, unspeciated aerosols (AP2.5), PMC, PM<sub>2.5</sub> and PM<sub>10</sub>.

To estimate these matrices, the "NMC" method is was used here to compute **B** for each variable by taking the differences between forecasts of different lengths valid at the same time (Parrish and Derber, 1992; Rabier et al., 1998). Differences between the 24-and 12-h WRF/CMAQ forecasts of 60 pairs (two pairs a-per day) of analysis variables valid at either 0000 or 1200 UTC over November 2016 are were used. The horizontal and vertical length scales of the correlation matrices are were estimated by using recursive filters (Purser et al., 2003). The vertical distribution of the background error

SDs is shown in Figure 3, which varies with height and species, is shown in Figure 3. The vertical profile of the background error SDs corresponds to the vertical concentration distribution. This means that higher concentrations tend to have larger background error SDs (e.g., CO and nitrate). These SDs exhibit a common reduction with as the height increases, especially at the top of the boundary layer. The horizontal correlation of the background error determines the propagation of observation information in this direction, while whereas the vertical correlation determines the vertical extension of such increments. For gaseous pollutants and most individual aerosol components, the horizontal length scales increasede with height, whereashile for the total particulate matter (i.e., PM<sub>2.5</sub>, and PM<sub>10</sub>), the scales increased with height in the boundary layer and decreased with height in the free troposphere. The ground-level scale generally spreads 40–45 km for all control variables on average. The vertical length scale of most species first increaseds first and then decreaseds with height, which may be related to the vertical mixing (Kahnert, 2008) and stack emissions at approximatelyabout 200 m height.

# 2.1.4 EnKF assimilation algorithm

In EnKF, the time-dependent uncertainties of the state variables are estimated using a Monte Carlo approach through an ensemble. Uncertainty can be propagated with using linear or nonlinear dynamic models (flow-dependent background error covariance) by simply implementing ensemble simulations. The EnSRF algorithm introduced by Bierman (1977) and Maybeck (1979) is was used to constrain pollution emissions in this study. EnSRF is a deterministic EnKF that obviates the need to perturb observations, which has a higher computational efficiency and a better performance (Sun et al., 2009). The perturbation of the prior emissions represents the uncertainty. We implemented additive emission adjustment methods, which are were calculated using the following function:

411 
$$X_i^b = X_0^b + \delta X_i^b, i = 1, 2, ..., N$$
 (7)

where b represents is the background (prior) state, i is the identifier of the perturbed

samples, and N is the ensemble size, which was set to 40 in consideration of considering thea trade\_off between the computational cost and inversion accuracy (Figure S1). In contrast to the estimation of parameters based on the augmentation of the conventional state vector (e.g. concentrations) with the parameter variables, X only comprises emissions in this study (similarly hereafter).  $\delta X_i^b$  represents is the randomly perturbed samples that are added to the prior emissions  $X_0^b$  to produce ensemble samples of the inputs  $X_i^b$ .  $\delta X_i^b$  is drawn from Gaussian distributions with a mean of zero and the standard deviation of the prior emission uncertainty in each grid. The state variables of the emissions include CO, SO<sub>2</sub>, NO<sub>x</sub>, primary PM<sub>2.5</sub> (PPM<sub>2.5</sub>) and PMC. We used variable localization to update the analysis, which means that the covariance among different state variables was not considered, and the emission of one species was only constrained only by with its corresponding air pollutant observation. This method has been widely used in chemical data assimilation systems to avoid spurious correlations among between species. (Ma et al., 2019; Miyazaki et al., 2012b).

After obtaining an ensemble of state vectors (prior emissions), ensemble runs of the CMAQ model are were conducted to propagate these errors in the model with each ensemble sample of state vectors. Combined with the observational vector y, the state vector is was updated by minimizing the analysis variance.

$$\overline{X^a} = \overline{X^b} + K(y - H\overline{X^b}) \tag{8}$$

$$\mathbf{K} = \mathbf{P}^{\mathbf{b}} \mathbf{H}^{T} (\mathbf{H} \mathbf{P}^{\mathbf{b}} \mathbf{H}^{T} + \mathbf{R})^{-1}$$
(9)

433 
$$\mathbf{P}^{b} = \frac{1}{N-1} \sum_{i=1}^{N} (\mathbf{X}_{i}^{b} - \overline{\mathbf{X}}^{b}) (\mathbf{X}_{i}^{b} - \overline{\mathbf{X}}^{b})^{T}$$
 (10)

$$\delta X_i^a = \delta X_i^b - \widetilde{K} H \delta X_i^b \tag{11}$$

While employing sequential assimilation and independent observations,  $\tilde{\mathbf{K}}$  is calculated as follows:

437 
$$\widetilde{\mathbf{K}} = \left(1 + \sqrt{\mathbf{R}/(\mathbf{H}\mathbf{P}^b\mathbf{H}^T + \mathbf{R})}\right)^{-1}\mathbf{K}$$
 (12)

where  $\overline{X}^b$  represents—is the mean of the ensemble samples; H is the observation operator that maps simulated concentrations from the model space to the observation space;  $y - H\overline{X}^b$  reflects the differences between the simulated and observed concentrations;  $P^b$  is the ensemble-estimated background (a priori) error covariance;  $P^bH^T$  contains the response of the uncertainty in the simulated concentrations to the uncertainty in emissions; K is the Kalman gain matrix of the ensemble mean depending on the  $P^b$  and observation error covariance R, representing the relative contributions to analysis; and  $\widetilde{K}$  is the Kalman gain matrix of the ensemble perturbation, which is used to calculate emission perturbations after inversions  $\delta X_i^a$ . The ensemble mean  $\overline{X}^a$  of the analyzed state is taken aswas considered the best estimate of the emissions.

With When large volumes of site observations that are recorded at a much higher resolution than the model grid spacing, manythere would be significant correlated or fully consistent model-data mismatch errors can appear in one cluster, resulting in excessive adjustments and deteriorated model performances (Houtekamer and Mitchell, 2001). To reduce the horizontal observation error correlations and the influence of representativeness errors, a "super-observation" approach combining multiple noisy observations located within the same grid and assimilation window is was developed based on optimal estimation theory (Miyazaki et al., 2012a). Previous studies have demonstrated the necessity of for data-thinning and dealiasing errors (Feng et al., 2020b; Zhang et al., 2009a). The super-observation  $y_{new}$ , super-observation error  $r_{new}$  and corresponding simulation  $x_{new,i}$  of the *i*th sample are calculated as follows:

$$\frac{1}{r_{new}^2} = \sum_{j=1}^m \frac{1}{r_j^2} \tag{13}$$

460 
$$y_{new} = \sum_{j=1}^{m} w_j y_j / \sum_{j=1}^{m} w_j$$
 (14)

461 
$$x_{new,i} = \sum_{j=1}^{m} w_j \, x_{ij} / \sum_{j=1}^{m} w_j$$
 (15)

where j is the identifier of m observations within a super-observation grid;  $r_j$  is the observational error of the actual jth observation  $y_j$ ;  $x_{ij}$  represents a is the simulated

concentration using the *i*th prior emission sample corresponding to the *j*th observation; 464 and  $w_j = \frac{1}{r_i^2}$  is the weighting factor. The super-observation error decreaseds as the 465 466 number of observations used within a super-observation increaseds. This method has 467 beenwas used in our previous inversions using surface-based (Feng et al., 2020b) and 468 satellite-based (Jiang et al., 2021) observations. 469 In this study, the DA window was set to 1-one day because the model need requires a 470 longer time to integrate the emission information into the concentration ensembles (Ma 471 et al., 2019). Due to the "super-observation" approach, only one assimilation is needed 472 in one assimilation window. In addition, due toowing to the complexity of hourly 473 emissions, it is very difficult to simulate hourly concentrations that can match the 474 observations well. Although a longer DA window could allow more observations 475 to constrain the emission change of one grid, the spurious correlation signals of EnKF 476 would attenuate the observation information overwith time (Bruhwiler et al., 2005; 477 Jiang et al., 2021). Kang et al. (2012) conducted OSSEs and demonstrated that due towing to the errors of transport errors and increased the spurious correlation, a longer 478 479 DA window (e.g., 3 weeks) would cause the analysis system to blur-out the essential 480 emission information-far away from the observation. Therefore, daily mean simulations 481 and observations wereare used in the EnSRF algorithm, and daily emissions wereare 482 optimized in this system. 483 EnKF is subject to spurious correlations due to because of the limited number of 484 ensembles when it is applied in high-dimensional atmospheric models, which can cause 485 rank deficiencies in the estimated background error covariance and filter divergence, 486 and further degrade analyses and forecasts (Wang et al., 2020). Covariance localization 487 is performed to reduce spurious correlations caused by the a finite ensemble size 488 (Houtekamer and Mitchell, 2001). Covariance localization preserves the meaningful 489 impact of observations on state variables within a certain distance (cutoff radius) but 490 limits the detrimental impact of observations on remote state variables. The localization

function of Gaspari and Cohn function (Gaspari and Cohn, 1999) is used in this system, which is a piecewise continuous fifth-order polynomial approximation of a normal distribution. The optimal localization scale is related to the ensemble size, assimilation window, dynamic system, and lifetime of a the chemical species in the atmosphere. CO, SO<sub>2</sub> and PM<sub>2.5</sub> are rather stable in the atmosphere, with a lifetime of more than one-1 day. According to the averaged wind speed (3.3 m/s, Table 4) and the length of the DA window, their the localization scales of CO, SO<sub>2</sub> and PM<sub>2.5</sub> are were set to 300 km. In addition, the localization scales of NO<sub>2</sub>, which is rather reactive and has, with a lifetime of approximately 10 hours in winter (de Foy et al., 2015), and PMC, which is mainly from local sources and has a short, its residence time in the atmosphere is also short due toowing to the rapid deposition rate (Clements et al., 2014; Clements et al., 2016; Hinds, 1982), Their localization scales arewere set to 150 km and 250 km, respectively.

#### 2.2 Prior emissions and uncertainties

The aAnthropogenic emissions over China were taken obtained from the 2016 Multiresolution Emission Inventory for China (MEIC 2016) (Zheng et al., 2018), while those over the other regions of East Asia were obtained from the mosaic Asian anthropogenic emission inventory (MIX) (Li et al., 2017). The spatial resolutions of both the MEIC and MIX inventories were both are  $0.25^{\circ} \times 0.25^{\circ}$ , and they are both downscaled to match the model grid spacing of 36 km. The spatial distributions of the CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC emissions are shown in Figure 12. The daily emission inventory, which was arithmetically averaged from the combined monthly emission inventory, was directly used in the EI subsystem and was employed as the prior emission of the first DA window in the EI subsystem (Figure 1). During the simulations, the daily emissions were further converted to hourly emissions. For aAll the species emitted from area sources, we were converted them to hourly emissions using a the same diurnal profile (Figure S2), and for the point source, we assumed that there was no diurnal change. MEIC 2012 was used as an alternative a priori over China to investigate the impact of different prior emissions on the optimized emissions. The Model of Emissions of Gases and Aerosols from Nature (MEGAN) (Guenther et al., 2012) was used to calculate timedependent biogenic emissions. It, which was also driven by the WRF model in this study. Biomass burning emissions were not included because they have little impact across China during the study period (Zhang et al., 2020).

During the inversion cycles, the inverted emissions of different members converge gradually gradually, and the ensemble-estimated error covariance matrix is arithmetically likely to be underestimated. To avoid this, considering the compensation of model errors and comparable emission uncertainties from one day to the next, we imposed the same uncertainty on emissions at each DA window. As mentioned above, the optimized emissions of the current DA window are were transferred to the next DA window as prior emissions. The technology-based emission inventory developed by Zhang et al. (2009b), basically using the same method as MEIC, showeds that the emissions of PMC and PPM<sub>2.5</sub> have had the largest uncertainties, followed by CO, and finally SO<sub>2</sub> and NO<sub>x</sub>. Therefore, the uncertainties of PMC, PPM<sub>2.5</sub>, CO, SO<sub>2</sub>, and NO<sub>x</sub> in this study are were set to as 40%, 40%, 30%, 25%, and 25%, respectively. However, previous studies have shown that the inversely estimated CO and PMC emissions could can exceed 100% higher than the bottom-up emissions (MEIC) in certain areas (Feng et al., 2020b; Ma et al., 2019). Therefore, aAccording to the extent of underestimation, we set an uncertainty of 100% for both the CO and PMC emissions at the beginning of the three DA windows to quickly converge the emissions. The mMean emission analysis is generally minimally sensitive to the uncertainty setting in our the assimilation cycle method (Feng et al., 2020; Gurney et al., 2004; Miyazaki et al., 2012a) because as the inversion errors of the current window could can be transferred to the next window for further optimization (Section- 4.3).

#### 2.3 Observation data and errors

520

521

522

523

524

525

526

527

528

529

530

531

532

533

534

535

536

537

538

539

540

541

542

543

544

545

546

547

548

Hourly averaged surface CO, SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>, PM<sub>2.5</sub> and PM<sub>10</sub> observations from 1504 national control air quality stations were assimilated into this system, which were obtained from the Ministry of Ecology and Environment of the People's Republic of China (http://106.37.208.233:20035/, last access: 25 June 2020). These sites are distributed over most of central and eastern China and become denser near metropolitan

areas (see Figure 2). Value-range and time-continuity checks were performed tTo ensure data quality, v-Value-range checks were mainly-performed to eliminate unrealistic or unrepresentative observations and other observations within the subjectively selected threshold range were assimilated (Table 2). In additionally, aA time-continuity check was performed to eliminate gross outliers and a sudden anomalies using thea function of  $max(|O(t) - O(t \pm 1)|) \le f(t)$ , where O(t) and  $O(t \pm 1)$  represent observations at time t and  $t \pm 1$ , respectively, and  $f(t) = T_a + T_b \times O_t$ . That This means that both-the concentration differences between time t and time t+1 and t-1should be less than f(t).  $T_b$  is was fixed to at 0.15, and the section of  $T_a$  is given in Table 2, which is was determined empirically according to the time series change of concentration at each site. It should be noted that, tTo avoid potential cross-correlations, we assimilated PM<sub>2.5</sub> and PMC. Additionally, in the EI subsystem, the observations within each city were averaged to thin reduce the data density, reduce the error correlation, and increase the spatial representation (Houtekamer and Mitchell, 2001; Houtekamer and Zhang, 2016). Finally, 336 city sites are—were available across the mainland of China, in which data from 311 cities data were selected for assimilation and the remaining 25 were selected for independent validation (Figure 2). In the IA subsystem, due owing to the small horizontal correlation scale (Figure 3), to obtain more extensive observation constraints, all site observations were assimilated to provide a good IC for the next emission inversion to obtain more extensive observation constraints.

549

550

551

552

553

554

555

556

557

558

559

560

561

562

563

564

565

566

567

568

569

570

571

573

574

575

576

The observation error covariance matrix (R) includes both the measurement and representation errors. The measurement error  $\varepsilon_0$  is defined as follows:

$$\varepsilon_0 = ermax + ermin \times \Pi_0 \tag{16}$$

where ermax is the base error, and  $\Pi_0$  denotes the observed concentration. These parameters for different species are listed in Table 2, which are and were determined according to Chen et al. (2019), Feng et al., (2018), and Jiang et al. (2013b).

The representative error depends on the model resolution and the characteristics of the

observation locations, which were calculated using the equations of Elbern et al. (2007), defined as follows:

$$\varepsilon_r = \gamma \varepsilon_0 \sqrt{\Delta l/L} \tag{17}$$

where  $\gamma$  is a tunable parameter (here,  $\gamma$ =0.5),  $\Delta l$  is the grid spacing (36 km), and L indicates is the radius (here, 3 km for simplification) of the influence area of an the observation. The total observation error (r) is was defined as follows:

$$r = \sqrt{\varepsilon_0^2 + \varepsilon_r^2} \tag{18}$$

Table 2. Parameters of quality control and measurement error

Parameter	CO mg m <sup>-3</sup>	SO <sub>2</sub> μg m <sup>-3</sup>	NO <sub>2</sub> μg m <sup>-3</sup>	O <sub>3</sub> μg m <sup>-3</sup>	PM <sub>2.5</sub> μg m <sup>-3</sup>	PMC μg m <sup>-3</sup>	
value-range	0.1-12	1-800	1-250	1-250	1-800	1-900	
time-continuity $(T_a)$	2.5	160	70	80	180	180	
ermax	0.05	1	1	1	1.5	1.5	
ermin	0.5%	0.5%	0.5%	0.5%	0.75%	0.75%	

# 3 Experimental design

RAPAS was conducted according to the procedure and settings described in Section- 2. December is one of the months with the most severe air pollution, while whereas July is one of the least polluted months in China. Therefore, this study mainly tested the performance of the RAPAS system in over these two months. For December, the IA subsystem was run from 26 November to 31, November 2016, with a 6-hour interval cycling assimilation to optimize ICs (ICDA). A better IC at 0000 UTC on 1 December 1 can could be obtained by a five 5-day high-frequency cycling assimilation and atmospheric mixing. Then the EI subsystem was then run for December 2016 with a one-1-day assimilation window to optimize emissions (EMDA). In For July, the system also operated identically to that of in the same way as for December. It needs to should

be noted that due to wing to the stronger atmospheric oxidation, the lifetime of NO<sub>2</sub> in July is was significantly shorter than that in December; thus, we adopted a smaller localization scale for NO<sub>2</sub> (80 km). Both assimilation experiments used the combined prior emission inventories of 2016, as described in Section-2.2, and the emission base year coincideds with the research stage. An Observing Systems Simulation Experiment (OSSEE) was conducted to evaluate the performance of the RAPAS system, which has been widely used in previous assimilation systems development (Daley, 1997). In the OSSE experiment, we used the MEIC 2016 inventory as a "true" emission, and reduced the "true" emission by 30% over the mainland of China as a prior emission. The simulations were simulated using the "true" emission and sampled according to the locations and times of the real observations were used as artificial observations. The observation errors are were the same as those in EMDA. To evaluate the IC improvements from the IA subsystem, an experiment without 3DVAR (NODA) is was conducted with the same meteorological fields and physical and chemistry parameterization settings as those of the ICDA. To evaluate the posterior emissions of the EI subsystem, two parallel forward modelling experiments are were performed for December 2016, namely,: a control experiment (CEP) with prior (MEIC 2016) emissions and a validation experiment (VEP) with posterior emissions. Both experiments used the same IC at 0000 UTC on December 01 generated through the IA subsystem. Similar to the above, tThe only differences between CEP and VEP are were emissions. Table 3 gives a summary of summarizes the different emission inversion experiments conducted in this study. To investigate the robustness of our system, 8-eight sensitivity tests (from EMS1 to EMS8; see Table 3) are were performed. These experiments are were all based on EMDA. In EMS1, rather than forward simulationed using the optimized emissions of the previous DA window in EMDA, the ICs of each DA window were first taken from the forward simulation with the prior emissions of the previous DA window, and then optimized using the 3DVAR EnSRF algorithm and the observations at the corresponding moment, as mentioned in Section-2.3. The objective of this experiment

597

598

599

600

601

602

603

604

605

606

607

608

609

610

611

612

613

614

615

616

617

618

619

620

621

622

623

624

is—was\_to investigate the advantages of the "two-step" calculation scheme in the EI subsystem—as introduced in Sect. 2.1. EMS2 useds MEIC 2012 as the original prior emission in China, aiming to investigate the impact of different prior inventories on the estimates of emissions. The Four other experiments, namely—(EMS3—6), aimed to test the impact of different prior uncertainty settings, in which, the prior uncertainties are—were reduced by -50% and -25%, and increased by 25% and 50%, respectively. EMS7 aimeds to evaluate the impact of observation errors on emission estimates, in which all the observation errors are magnified twice. The last-EMS8 experiment aimeds to evaluate the impact of IC optimization of the first window on emission estimates, in which the ICs were taken from a five5-day spin-up simulation. Eight forward modelling experiments (VEP1, VEP2, ..., VEP8) were also performed with the posterior emissions of EMS1 to EMS8 to evaluate their performances, respectively.

Table 3. Emission inversion and sensitivity experiments conducted in this study

Exp. Type	Exp. Name	Period	IC of the first DA Window	ICs of the subsequent DA window	Emission
Assimilation	EMDA	131 December	0000 UTC on December 1, taken from ICDA	Forecast with posterior emissions in the previous window	MEIC 2016 for December (the first DA window), optimized emissions of the previous window (other DA windows)
	OSSE	1 <u>—</u> 31 December	The sSame as EMDA	The sSame as EMDA	The sSame as EMDA but with a decrease of 30% for CO, SO <sub>2</sub> , NO <sub>x</sub> , PPM <sub>2.5</sub> , and PMC
	EMS1	1_31 December	The sSame as EMDA	Optimized using the EnSRF DA methodForecast with prior emissions in the previous window and 3DVAR assimilation	The sSame as EMD
	EMS2	1 <u>-</u> 31 December	The sSame as EMDA	The sSame as EMDA	The sSame as EMDA but for EMIC 2012
Sensitivity	EMS3-6	131 December	The sSame as EMDA	The sSame as EMDA	The sSame as EMDA but with a $\pm 25\%$ or $\pm 50\%$ of default uncertainty
	EMS7	1 <u>-</u> 31 December	The sSame as EMDA	The sSame as EMDA	The sSame as EMDA but with a +100% or default observation errors
	EMS8	1 <u>-</u> 31 December	0000 UTC on December 1, taken from ICNO	The sSame as EMDA	The sSame as EMD

# 4 Results

658

659

660

661

662

663

664

665

666

667

668

669

670

671

672

673

674

675

676

677

678

679

680

681

682

683

684

685

#### 4.1 Evaluations

#### 4.1.1 Simulated meteorological fields

In the RAPAS system, the inversion approach attributes all the biases between the simulated and observed concentrations to the emissions. The mMeteorological fields dominate the physical and chemical processes of the air pollutants in the atmosphere, and thus their simulation accuracy would significantly affect the estimates of emissions in this study. To quantitatively evaluate the performance of the WRF simulations, the mean bias (BIAS), root mean square error (RMSE), and correlation coefficient (CORR) were calculated against the surface meteorological observations measured at 400 stations and the planetary boundary layer height (PBLH) was calculated using the sounding data at 92 sites. The sSurface observations were obtained from the National Climate Data Center(NCDC) integrated surface database (http://www.ncdc.noaa.gov/oa/ncdc.html, last access: 25 October 2021), and the sounding data were obtained from the website of the University of Wyoming (http://weather.uwyo.edu/upperair/sounding.html, last access: 10 March 2022). The sounding data are inhad a 12 hours interval. The observed PBLH were was calculated using the sound data through via the bulk Richardson number method (Richardson et al., 2013). The spatial distribution of the meteorological stations is shown in Figure 2. The simulated temperature at 2 m (T2), relative humidity at 2 m (RH2), wind speed at 10 m (WS10), and PBLH from 26 November to 31 December 2016 are-were evaluated against the observations. Table 4 summarizes the statistical results of the evaluations of the simulated meteorological parameters. Overall, the T2, RH2 and PBLH are were slightly underestimated, with biases of -0.1 °C, -3.8%, and -41.1 m, respectively. The CORRs are were approximately 0.98 for T2, 0.94 for RH2, and 0.90 for PBLH, showing good consistency between the observations and simulations. The WS10 is was overestimated, with a bias of 0.7 m/s and an RMSE of 0.8 m/s, but is were better than the simulations from many other previous studies (Chen et al., 2016; Jiang et al., 2012a;

Jiang et al., 2012b). Therefore, the WRF can generally reproduce the meteorological conditions sufficiently in terms of their temporal variation and magnitude over China, which is adequate for our inversion estimation.

**Table 4**. Statistics comparing the simulated and observed 10-m wind speed (WS10), 2-m temperature (T2), and 2-m relative humidity (RH2), and planetary boundary layer height (PBLH).

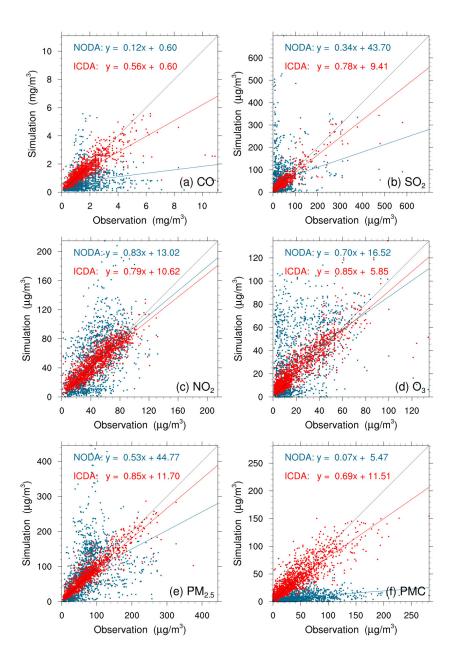
Variable Met.	No. of sites	Mean Obs.	Mean Sim.	BIAS	RMSE	CORR
WS10 (m/s)	400	2.6	3.3	0.7	0.8	0.72
T2 (°C)	400	2.9	2.8	-0.1	0.7	0.98
RH2 (%)	400	66.3	62.6	-3.8	5.2	0.94
PBLH (m)	92	267.5	226.4	-41.1	50.4	0.90

<sup>\*</sup> BIAS, mean bias; RMSE, root mean square error; CORR, correlation coefficient

# 4.1.2 Initial fields conditions

Figure 4 shows the an evaluations of the analyzed concentrations of the 6-six species against surface observations. For comparison, the evaluations of the simulations without 3DVAR (NODA) are also shown in Figure 4. The simulations of the NODA experiment (red dots) are scattered on both sides of a the central line, as large systematic biases remain across many measurement sites. Conversely, the ICDA experiment (blue dots) showed aws much better agreement with the observations than those from NODA. The statistics show that there are large systematic biases in the NODA simulations, with large RMSEs and small CORRs for all species, especially particularly for CO and PMC. After the assimilation of surface observations, the RMSE of CO decreaseds to 0.7 mg m<sup>-3</sup>, and those of SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>, PM<sub>2.5a</sub> and PMC decrease to 22.0, 12.0, 9.6, 20.5<sub>a</sub> and 19.6 μg m<sup>-3</sup>, respectively, with respective reductions rates of 50.0%, 73.1%, 61.0%, 64.7%, 69.5%, and 60.8% compared to the onesthose of the NODA (Table 5). The CORRs of ICDA increased by 290.0%, 291.3%, 55.4%, 87.2%, 130.0%<sub>a</sub> and 214.8% to 0.78, 0.90, 0.87, 0.88, 0.92<sub>a</sub> and 0.85, respectively. These statistics indicate that the

ICs of the ground level <u>improved have been</u> significantly <u>improved</u>. However, <u>due</u> <u>owing</u> to the lack of observations, we still do not know the simulation bias in the upper-middle boundary layer. Although concentrations at high altitudes can be constrained by ground-based observations through vertical correlations, the effect is limited; <u>therefore</u>, so the bias <u>is stillremains</u> non-negligible.



**Figure 4.** Scatter plots of simulated versus observed (a) CO, (b) SO<sub>2</sub>, (C) NO<sub>2</sub>, (d) O<sub>3</sub>, (e) PM<sub>2.5</sub> and (f) PMC mass concentrations at 0000 UTC on December 1 initializations from the background (red) and analysis (blue) fields.

**Table 5**. Comparisons of the surface CO, SO<sub>2</sub>, NO<sub>2</sub>, O<sub>3</sub>, PM<sub>2.5</sub> and PMC mass concentrations from the control and assimilation experiment against observations aggregated over all analysis times. CO unit: mg m<sup>-3</sup>; others units: μg m<sup>-3</sup>.

Species	Exp. Name	Mean Obs.	Mean Sim.	BIAS	RMSE	CORR
СО	NODA	1.5	0.8	-0.7	1.4	0.20
	ICDA	1.5	1.5	-0.1	0.7	0.78
$\mathrm{SO}_2$	NODA	36.3	56.0	19.7	81.7	0.23
	ICDA	30.3	37.8	1.5	22.0	0.90
NO <sub>2</sub>	NODA	45.8	51.1	5.3	30.8	0.56
	ICDA		47.0	1.1	12.0	0.87
$O_3$	NODA	20.5	30.8	10.4	27.2	0.47
	ICDA	20.3	23.3	2.8	9.6	0.88
PM <sub>2.5</sub>	NODA	70.9	82.2	11.3	67.3	0.40
F 1V12.5	ICDA		71.8	0.9	20.5	0.92
PMC	NODA	43.5	8.5	-35.0	50.0	0.27
1 1/10	ICDA		41.6	-1.9	19.6	0.85

\* BIAS, mean bias; RMSE, root mean square error; CORR, correlation coefficient

#### 4.1.3 Posterior emissions

Due Owing to the mismatched spatial scales, it is difficult to directly evaluate the optimized emissions against observations. Generally, we indirectly validated them the optimized emissions by comparing the forward simulated concentrations using the posterior emissions against atmospheric measurements (e.g., Jiang et al., (2014;), Jin et al., (2018); and Peters et al., (2007)). Figure 5 shows the spatial distributions of the mean biases between the simulated gaseous pollutants simulated using prior and posterior emissions and assimilated observations. In the CEPs, for each species, the distribution of biases is was similar to the increments in background fields constrained through 3DVAR, as shown in Figure S3. For example, almost all sites have had large negative biases for CO, while for SO<sub>2</sub> and NO<sub>2</sub>, positive biases are were mainly distributed over the North China Plain (NCP), Yangtze River Delta (YRD), Sichuan Basin (SCB), and Central China, and negative biases are were distributed over the rest

of theremaining areas. After constraining with observations, the biases of all the 3three gaseous air pollutants are were significantly reduced in at most sites. For CO, the biases at 62% of the sites decreased to absolute values less than 0.2 mg  $m^{-3}$ , and for SO<sub>2</sub> and NO<sub>2</sub>, the biases at 52% and 47% of the sites were within  $\pm 4 \mu g \text{ m}^{-3}$ . However, large negative biases are were still observed in part of western Chinas, indicating that the uncertainties of the posterior emissions are still large in western China, which may be attributed to the large biases in prior emissions and to the relatively limited observations. Overall, the statistics show that there are different levels of improvements at 92%, 85% and 85% of the total 311 assimilation sites of 92%, 85%, and 85% for CO, SO<sub>2</sub>, and NO<sub>2</sub>, respectively. The small amount number of sites with worse performance may be related to the over-adjusted emissions by EI or contradictory adjustments caused by opposite biases in adjacent areas. Table 6 lists the statistical results of the evaluations averaged over the whole mainland of China. For CO, the mean bias is-was -0.8 mg m<sup>-3</sup> with the prior emissions, while it substantially reduceds to -0.1 mg m<sup>-3</sup> (with a reduction rate of 89.6%) when simulating with the posterior emissions. Additionally, the RMSE decreases decreased by 48.1% from 1.08 to 0.56 mg m<sup>-3</sup>, and the CORR increased by 76.1% from 0.46 to 0.81. For SO<sub>2</sub> and NO<sub>2</sub>, the regional mean biases slightly increased as the positive/negative biases among different sites might be offset. However, the RMSEs decreased to 17.7 and 12.3 μg m<sup>-3</sup>, respectively, which are were 58.3% and 50.8% lower than those of CEPs, and the CORRs increased by 125.6% and 35.4%, both reaching up to 0.88, indicating that EI has significantly improved the  $NO_x$  and  $SO_2$  emission estimates.

734

735

736

737

738

739

740

741

742

743

744

745

746

747

748

749

750

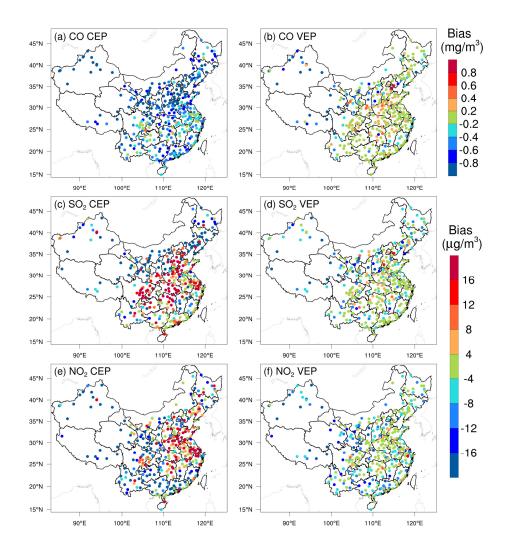
751

752

753

754

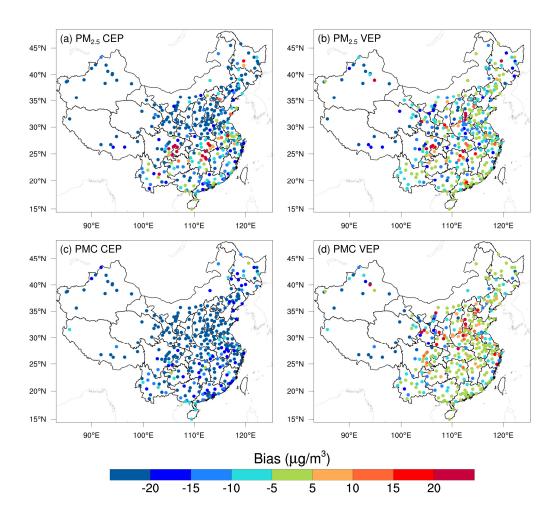
755



**Figure 5**. Spatial distribution of the BIAS of the simulated (a, b) CO, (c, d)  $SO_{2a}$  and (e, f)  $NO_2$  with prior (left, CEP) and posterior (right, VEP) emissions. CO unit: mg m<sup>-3</sup>;  $SO_2$  and  $NO_2$  units:  $\mu$ g m<sup>-3</sup>.

Figure 6 shows the spatial distributions of the mean biases of simulated PM<sub>2.5</sub> and PMC evaluated against the assimilated observations. Similarly, the CEP simulations do-did not perform well. There are were widespread underestimations across the country, with mean biases of -24.0 and -32.4  $\mu g$  m<sup>-3</sup>. After data assimilation, the performance of the VEP simulations is significantly improved. The biases decreased by 72.1% and 90.4% to -6.7 and -3.1  $\mu g$  m<sup>-3</sup>, the RMSEs decreased by 41.2% and 40.7% to 29.6 and 24.6  $\mu g$  m<sup>-3</sup>, and the CORRs increased by 35.9% and 176.0% to 0.87 and 0.69 for PM<sub>2.5</sub> and PMC, respectively. Overall, 89.6% and 97.2% of the assimilation sites are were

improved for PM<sub>2.5</sub> and PMC, respectively. However, compared with the results of for the 3-three gaseous pollutants, there are were sites with large biases scattered throughout the whole entire domain. Besides In addition to the potential over-adjusted or contradictory adjustments of emissions as in the 3-three gas species, It-the sites with large biases may also be related to the complex precursors and complex homogeneous and heterogeneous chemical reactions and transformation processes of secondary PM<sub>2.5</sub>, and the fact that we do did not simulate the time variation of dust blowing caused by wind speed for PMC due owing to the lack of land cover data that is compatible with the CMAQ dust module and agricultural activityies data to identify dust source regions.



**Figure 6**. Same as in Figure 5 but for PM<sub>2.5</sub> and PMC.

Figures 7 and Figure 8 show the spatial distributions of the biases calculated against the independent observations for the 5-five species. With posterior emissions, the

decreasing ratios of RMSEs ranged from 26.7%—to 42.0%, and the CORRs increased by 13.7—59.0% to 0.62—0.87. Overall, the biases at the independent sites are similar or slightly worse than those at the assimilated sites, which is reasonable since as the closer the independent sites are to the assimilated site—the independent sites are, the more constraints of observation information can be obtained, and the more significant the improvements in the optimized state variables of the model—are more significant. For example, generally, the transmission distance of NO<sub>2</sub> is relatively short, and remote cities with small emission correlations to the cities with assimilated observations are relatively less constrained, resulting in only a 26.7% decrease in the RMSE.

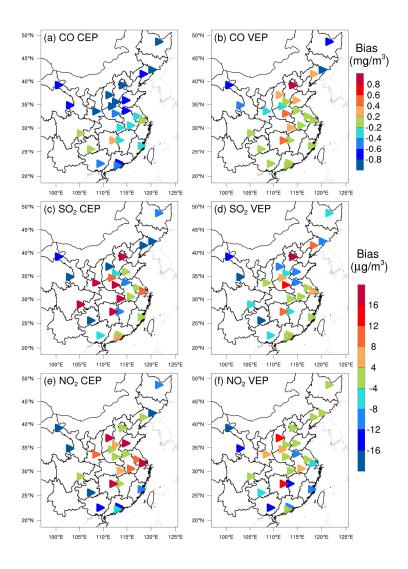
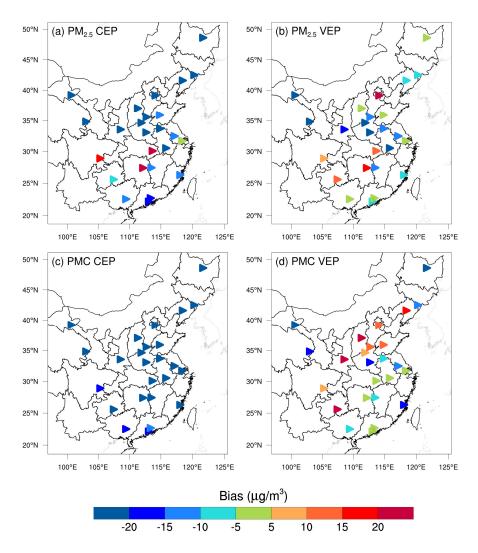


Figure 7. As in Figure 5 but for the independent validation.



**Figure 8**. As in Figure 6 but for the independent validation.

Comparing our resultsed with those of the previous studies, Tang et al. (2013) conducted inversion of inverted CO emissions over Beijing and the surrounding areas, and obtained comparable the improvements (Table 6) in the RMSE (37—48% vs. 30—51%) and the CORR (both studies ~ 0.81) are comparable; however, but the biases here could decrease we decreased the biases by 90—97%, which is much greater than their 48—64% reductions. Additionally, Chen et al. (2019) showed that the RMSE of simulated SO<sub>2</sub> with updated SO<sub>2</sub> emissions decreased by 4.2—52.2% for different regions, and the CORR only increased to 0.69 at most. These improvements is are relatively smaller than our results those obtained in this study, which may be due to the insufficient adjustment of emissions caused by the underestimated ensemble spread

through the inflation method. The better performance in this study may be related to our inversion process, which causes that makes the optimized emissions of the current DA window to propagate to the next DA window for further correction.

**Table 6**. Statistics comparing the pollution concentrations from the simulations with prior (CEP) and posterior (VEP) emissions against assimilated and independent observations, respectively. CO unit: mg m<sup>-3</sup>; others units: μg m<sup>-3</sup>.

Species	Mean	Mean Sim.		BIAS		RMSE		CORR	
	Obs.	CEP	VEP	CEP	VEP	CEP	VEP	CEP	VEP
Against assimilated observations									
CO	1.43	0.66	1.36	-0.77	-0.08	1.08	0.56	0.46	0.81
$SO_2$	32.5	34.4	28.4	1.9	-4.1	42.4	17.7	0.39	0.88
$NO_2$	43.8	40.8	39.0	-2.9	-4.8	25.0	12.3	0.65	0.88
PM <sub>2.5</sub>	77.0	53.1	70.3	-24.0	-6.7	50.3	29.6	0.64	0.87
PMC	40.5	8.1	37.5	-32.4	-3.1	41.5	24.6	0.25	0.69
Against independent observations									
CO	1.54	0.79	1.52	-0.75	-0.02	1.15	0.72	0.59	0.82
$SO_2$	40.6	39.2	37.3	-1.3	-3.2	44.3	27.2	0.57	0.87
$NO_2$	50.2	50.0	47.5	-0.3	-2.7	21.7	15.9	0.73	0.83
PM <sub>2.5</sub>	91.5	64.6	84.1	-26.9	-7.4	64.1	37.2	0.62	0.87
PMC	42.0	9.2	40.4	-32.8	-1.6	39.3	26.6	0.39	0.62

<sup>\*</sup> BIAS, mean bias; RMSE, root mean square error; CORR, correlation coefficient

### 4.1.4 Uncertainty reduction

The uncertainty reduction rate (UR) is another an important quantity to evaluate the performance of RAPAS and the effectiveness of <u>in situ</u> observations in this system (Chevallier et al., 2007; Jiang et al., 2021; Takagi et al., 2011). Following Jiang et al. (2021), the UR is was calculated as

817 
$$UR = \left(1 - \frac{\sigma_{posterior}}{\sigma_{prior}}\right) \times 100 \tag{19}$$

where  $\sigma_{posterior}$  and  $\sigma_{prior}$  are the posterior and prior uncertainties, respectively, which were calculated using the standard deviations of the prior and posterior perturbations (Text S23). Figure 9 shows the URs averaged in each province and the whole mainland China. The URs variedy with species as they are closely related to the

magnitude settings of prior uncertainties (Jiang et al., 2021).7 The URs of PPM2.5 and PMC were the most effective while the UR of NO<sub>x</sub> emissions was the lowest among the 5 species of emissions, the uncertainties of the PPM2.5 and PMC are greatly reduced, while the UR of NO<sub>x</sub> emission is lowest, that is because the URs are closely related to the magnitude settings of prior uncertainties (Jiang et al., 2021). For the whole mainland China overall, the uncertainties are were reduced by 44.4%, 45.0%, 34.3%, 51.8%, and 56.1% for CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM2.5, and PMC, respectively. For one species, it URs also varieds across provinces. The URs are usually related to observation coverage, which means that the more observation constraints there are, the more the URs decrease. Additionally, the URs may also be related to emission distributions. Generally, the URs are were more significant in the provinces where the observations and emissions are were both relatively concentrated (e.g., Tibet), while they are were much lower in where the emissions are were scattered or relatively uniform, but the observations are were only in large cities, even though if there are were many more observations than in other provinces.

		Uncertai	nty reduc	ction (%)		
Mainland-	44.4	45.0	34.3	51.8	56.1	
Shanghai-	16.9	16.7	20.8	24.7	18.5	
Jiangsu-	17.7	25.3	29.3	34.1	52.3	
Zhejiang-	24.7	13.3	17.9	42.4	31.4	70
Anhui-	20.1	52.7	39.1	58.1	40.9	
Shandong-	32.1	30.0	20.3	53.7	26.7	65
Beijing-	28.2	6.2	37.0	43.3	31.4	
Tianjin	20.0	7.0	21.4	41.3	17.8	60
Hebei-	29.5	40.2	28.8	56.0	30.3	
Shanxi-	38.4	37.9	22.5	55.3	35.0	-55
Neimenggu-	30.1	45.8	40.4	37.6	52.8	F-0
Henan-	27.4	16.1	21.9	53.7	30.8	50
Hunan-	36.0	27.7	34.4	16.9	41.6	45
Hubei-	30.8	16.6	26.0	46.4	46.5	45
Jiangxi-	20.9	28.4	29.4	47.0	46.7	40
Guangdong-	31.2	14.9	41.1	53.1	46.4	40
Guangxi-	22.6	13.9	42.5	48.1	55.2	35
Fujian-	9.9	8.1	31.9	31.6	49.2	
Hainan	0.6	0.5	4.5	0.7	23.3	-30
Liaoning-	35.6	34.6	19.0	33.9	54.0	
Heilongjiang-	29.9	27.7	17.4	42.0	65.2	-25
Jilin-	27.9	44.5	18.7	42.0	42.8	
Shaanxi-	41.3	13.2	29.8	47.9	43.1	-20
Gansu-	24.8	36.1	33.7	46.3	56.4	
Xinjiang-	38.3	27.9	20.2	46.3	66.5	15
Qinghai-	53.9	25.8	27.3	46.0	57.9	
Ningxia-	47.0	36.6	17.6	38.0	30.1	10
Sichuan	29.4	25.0	39.5	61.1	46.5	
Chongqing-	5.7	8.2	8.8	12.7	13.8	-5
Guizhou <sup>-</sup>	14.4	16.4	26.6	40.3	38.2	
Yunnan-	38.3	29.9	31.4	40.1	55.9	-0
Tibet <sup>.</sup>	30.2	0.5	52.8	67.3	73.2	
	CO	$SO_2$	NO <sub>X</sub>	$PPM_{2.5}$	PMC	

**Figure 9**. Time-averaged posterior emission uncertainty reduction (%) indicated by the standard deviation reduction of total emissions per province calculated by prior and posterior ensembles.

# 4.1.5 Evaluation using chi-squared statistics

To diagnose the performance of the EnKF analysis, the chi-squared ( $\chi^2$ ) statistics was were calculated, which is are generally used to test whether the prior ensemble mean

RMSE with respect to the observations is consistent with the prior "total spread" (square root of the sum of ensemble variance and observation error variance). Following Zhang et al. (2015), for the *t*th window,  $\chi^2$  is defined as:

847 
$$\chi_t^2 = (\mathbf{y} - \mathbf{H}\overline{\mathbf{X}^b})^T (\mathbf{H}\mathbf{P}^b\mathbf{H}^T + \mathbf{R})^{-1} (\mathbf{y} - \mathbf{H}\overline{\mathbf{X}^b})$$
 (20)

Figure 10 shows the time series of the relative changes between the prior and posterior emissions and the  $\chi^2$  statistics. There are—were relatively large adjustments of—in emissions in the first three windows, especially for the PMC. SubsequentlyAfter that, the optimality of—the five species reacheds a more optimal state with successive emission inversion cycles. The  $\chi^2$  statistics shows—showeda similar variation characteristics with—as the daily changes in the emissions. The  $\chi^2$  value is—was slightly greater than 1, indicating that the uncertainties from the error covariance statistics do did not fully account for the error in the ensemble simulations. A similar result was reported by situation also appeared in Chen et al. (2019). Further investigations should be conducted to generate larger spreads by accounting for the influence of model errors. Since As we imposed thea same uncertainty of prior emissions at each DA window to partially compensate for the influence of model errors,  $\chi^2$  statistics showed small fluctuations, indicating that the system updates—updated emissions consistently and stably.

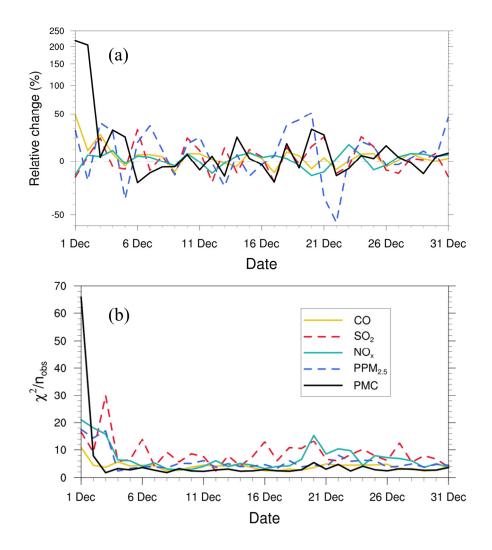
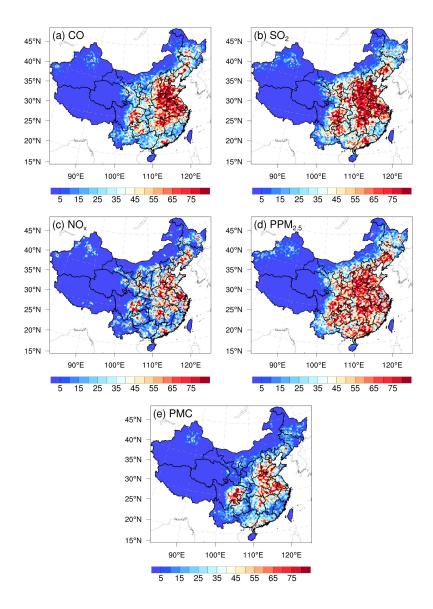


Figure 10. Relative changes (a) in a posterioriposterior emission estimates of CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5a</sub> and PMC\_-and  $\chi^2$  statistics (b) of these state vectors in each window.

### 4.1.6 Evaluation using OSSE

Figure 11 shows the spatial distribution of the error reduction in the posterior emissions of the five species. It can be found that aAfter inversion, in most areas, the emission errors can be were reduced by more than 80%, especially in the central and eastern regions with dense observation sites, while in remote areas far away from cities, due to the sparse observation sites, the emission errors are were still not well adjusted. Overall, the error reduction rates of CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC are were 78.4%, 86.1%, 78.8%, 77.6%, and 72.0%, respectively, indicating that with the ground in situin situ observations in China, RAPAS can significantly reduce emission errors and thus has showed good performance in emission estimates.



**Figure 11** Spatial distribution of the error reduction (%) of posterior emissions in the OSSE.

#### 4.2 Inverted emissions

Figure 12 shows the spatial distribution of the temporally averaged prior and posterior emissions and their differences of thein emissions in December 2016. It should be noted that the emissions outside China were masked; assince the observation sites are all withinwere limited to China in this study, there is was little a slight change in the emissions outside China. Higher emissions are were mainly concentrated in central and eastern China, especially in the NCP, YRD, and PRD, and lower emissions occurred across Northwest and Southern China. Compared with the prior emissions, posterior

CO emissions are were considerably increased across most areas of mainland China, especially in northern China, with an overall increase of 129%. A nNotable underestimation of the prior emissions is was also confirmed by previous inversion estimations (Feng et al., 2020b; Tang et al., 2013; Wu et al., 2020) and model evaluations (Kong et al., 2019b) in previous studies. For SO<sub>2</sub>, the emissions increases <u>increased</u> mainly <del>occur</del> in Northeast China, Shanxi, Ningxia, Gansu, Fujian, Jiangxi, and Yunnan provinces. In SCB, Central China, YRD, and part of the NCP, the emissions are were significantly reduced. For The national total, the SO<sub>2</sub> emissions is increased by 20%. For NO<sub>x</sub>, although the increment of national total emissions is was small (approximately only about 5%), there were large deviations still exist on regional scale. Obviously, tThe emissions in the NCP and YRD are were reduced, while whereas in the other regions, the emissions of in most cities in other regions are increased. The changes in the emission of PPM<sub>2.5</sub> emission are were similar to those of SO<sub>2</sub>. Compared with the prior emissions, the posterior PPM<sub>2.5</sub> emissions are decreased over central China, SCB, and YRD, while whereas the onesthose in southern and northern China are increased, especially in Shanxi, Shaanxi, Gansu, and southern Hebei provinces. Overall, the relative increase is was 95%. For PMC, the posterior emissions are were increased over the wholeall of mainland China, with national mean relative increase exceeding 1000%. Larger emission increments mainly occurred in the areas where have with significant anthropogenic emissions of CO and PPM<sub>2.5</sub>, indicating that the large underestimations of PMC emissions in the prior inventory may be mainly attributed to the underestimations of anthropogenic activities. In addition, tThe absence of natural dust is another reason, as the wind-blown dust scheme was not applied in this study. Overall, PM10 emissions (PPM<sub>2.5</sub>+PMC) increased by 318%. If we assume that all the increments in PM<sub>10</sub> emissions is allare from natural dust, that means the contribution of natural dust accounted for 75% of total PM<sub>10</sub> emissions, which is consistent with the source apportionment of PM<sub>10</sub> of 75% in Changsha in Central China (Li et al., 2010). Large PMC emission increments are were also found reported by in Ma et al. (2019).

887

888

889

890

891

892

893

894

895

896

897

898

899

900

901

902

903

904

905

906

907

908

909

910

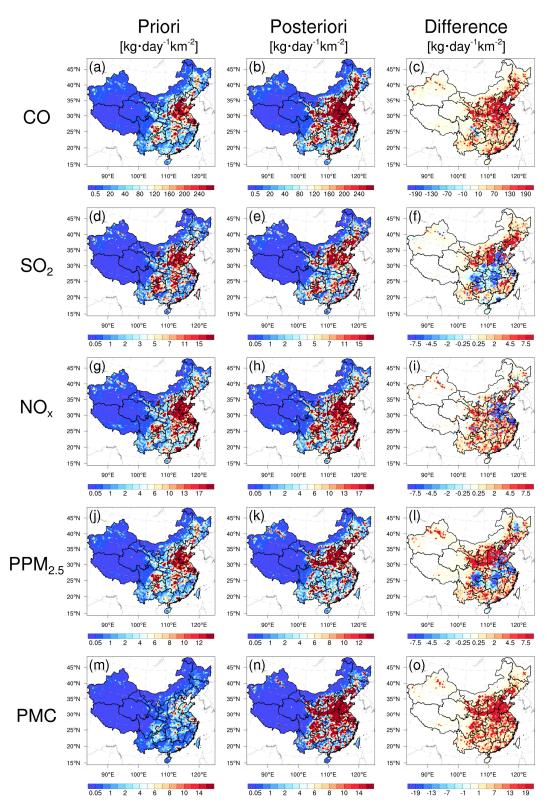
911

912

913

914

Detailed estimations of posterior emissions and relative changes compared to prior emissions in each province and the whole mainland China is are given in Table S1. The evaluation results for July showed that the emission uncertainty ean-could still be significantly reduced, and the performance of the system in July is was comparable to that in December (Table S2). Additionally, the seasonal variation of in emissions ean bewas well reflected (Figures S4 and S5), which means that our system ean performed well at different times of the year. Note that the differences, excluding PMC, between the prior and posterior emissions mainly reflect the deficiencies of the prior emissions because as the times of the prior emissions and the observations are completely were consistent in this study.



**Figure 12**. Spatial distribution of the time-averaged prior emissions (left column, MEIC 2016), posterior emissions (middle column), and differences (right column, posterior minus prior).

### 4.3 Sensitivity tests

930

931

932

933

934

935

936

937

938

939

940

941

942

943

944

945

946

947

948

949

950

951

952

953

954

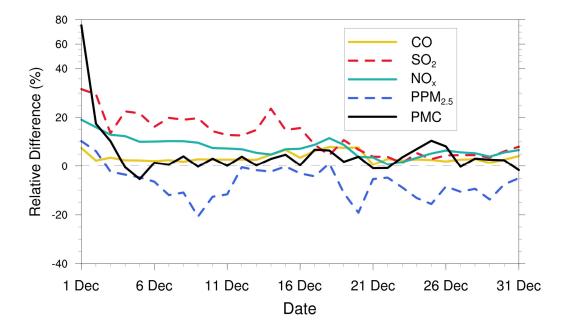
955

956

957

### 4.3.1 Impact of prior inventories

Various prior inventories have great shown considerable differences in space allocation and emission magnitudes. Inversion results can be sensitive to a priori emissions if the observations is are insufficient (Gurney et al., 2004; He et al., 2018). MEIC 2012 is was used as an alternative a priori in EMS2 to investigate the impact of different prior emissions on the posterior emissions. Figure 13 shows the time series of the relative differences in the daily posterior emissions of the five species between the EMDA (base) and EMS2 experiments. Overall, the differences between the two posterior emissions gradually decreased over time. At the beginning, the differences in the CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC between the two inventories (i.e., MEIC 2012 vs. MEIC 2016) are were 17.5%, 114.5%, 30.8%, 46.0%, and 72.0%, respectively, while during the last ten days, the differences of the two posterior emissions have decreased compared to 2.5%, 4.5%, 4.5%, -8.9%, and 3.0%, respectively in the last ten days. In addition, it also could be found that the species that has with larger emission differences at the beginning take took a longer time (namely i.e. more DA steps) to achieve convergence. The quick convergence of PMC emissions is was attributed to the large prior uncertainty of 100% used in the first 3-three DA windows. Different from In contrast to the other species, there are were significant negative deviations of in PPM<sub>2.5</sub> emissions between the two experiments. That This may be due to the positive deviations in the precursors of PM<sub>2.5</sub> (i.e.,  $SO_2$  and  $NO_x$ ), which will-lead to a larger amount of secondary production. To balance the total PM<sub>2.5</sub> concentration, tThe PPM<sub>2.5</sub> emissions will be reduced to balance the total PM<sub>2.5</sub>. We compared the PM<sub>2.5</sub> concentrations simulated by the two optimized inventories and find found that they are were almost the same (Figure S6). Overall, this indicates that the observations in China is were sufficient in to inferring the emissions, and that our system is ratherwas robust. Meanwhile, it also suggests that the monthly posterior emissions shown in Section- 4.2 are were still underestimated to a certain extent.



**Figure 13**. Relative differences in CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC emissions (%, the ratio of absolute difference to EMDA) between the EMDA and EMS2 experiments.

### 4.3.2 Impact of prior uncertainties settings

The uncertainty of prior emissions determines how closely the analysis is weighted towards the background and observations; however, but information about prior uncertainties is generally not readily available. To evaluate the possible influence of prior uncertainties on the optimized emissions, we increased/reduced the uncertainties after 3-three days of cycling, namely starting at 0000 UTC, 3 December, by 25% and 50% in EMS3 (-50%), EMS4 (-25%), EMS5 (+25%), and EMS6 (+50%), respectively. Table 7 summarizes the emission changes with different prior uncertainties uncertainty settings in the EMS3—6 experiments. To better understand the response of the system to the emission uncertainty settings, Figure 14 shows illustrates the time series of SO<sub>2</sub> emission changes, the Chi-square statistics, and the RMSEs of simulated SO<sub>2</sub> with emissions updated in the EMDA and EMS3—6 experiments over the YRD and NCP (Figure 2). Compared with the EMDA, when the uncertainties are decreased (increased), the emissions of the 5-five species decreased (increased) accordingly. That This is because the posterior emissions of the 5-five species are—were larger than the prior

emissions, and, as shown in Figure 14a-d, larger uncertainty will lead to a faster convergence, resulting in larger posterior emissions. It also could can also be seen found from Figure 14 that a faster convergence will indeed reduce the RMSE of the simulated concentration with the posterior emissions in the early stage of the experiment; however, but in the later stage of the experiment, there are were no significant differences for in the RMSE and Chi-square statistics among the different experiments. However, the day-to-day changes in emissions can also cause slight fluctuations. In addition, it shows that when greater uncertainties are set, the day-to-day changes in emissions are also-more drastic, resulting in a larger RMSE, as shown in the NCP. Moreover, those the significant day-to-day variations of in the estimated emissions may not be in line with the actual situation. Due Owing to the spatial-temporal inhomogeneity of emissions, the differences of in Chi-square statistics between the YRD and NCP show that it may be necessary to apply different a priori uncertainties according to different regions (Chen et al., 2019). Therefore, when using an EnKF system for emission estimation, error setting must be carefully executedwe have to be very careful about the setting of these errors. Overall, the uncertainties chosen in EMDA aim to minimize the deviation of the concentration fields and maintain the stability of the inversion.

**Table 7**. Relative differences in CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub> and PMC emissions (%, the ratio of absolute difference to EMDA) between the EMDA and EMS3-6 experiments.

Species	EMS3	EMS4	EMS5	EMS6
СО	-8.6	-4	3	5.2
$SO_2$	-14	-5.7	3.6	6.8
$NO_x$	-6.5	-3	2.8	4.5
PPM <sub>2.5</sub>	-16.5	-7.8	4.6	8.7
PMC	-18.5	-8.2	7.3	13.1

977

978

979

980

981

982

983

984

985

986

987

988

989

990

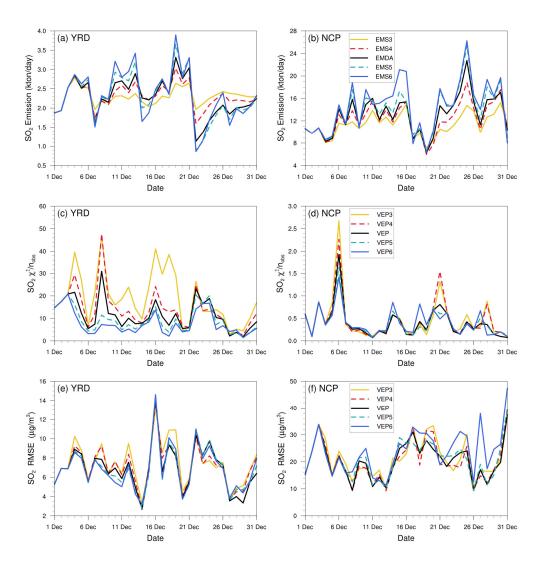
991

992

993

994

995



**Figure 14**. Time-series of SO<sub>2</sub> emission changes, the Chi-square statistics, and the RMSE of simulated SO<sub>2</sub> with updated SO<sub>2</sub> emissions in the EMDA and EMS3-6 experiments over the Yangtze River Delta (YRD) and North China Plain (NCP).

## 4.3.3 Impact of observation error settings

Observation errors are aAnother factor that determines the relative weights of the observations and background in the analysis is observation errors. A proper estimate of the observation error is also important in regard to thefor filter performance; buhowever, to observation errors are generally not provided with the datasets. The observation error is usually set to a fixed value (Ma et al., 2019), a specific proportion of the observation value (Tang et al., 2013), or the value calculated by combining measurement error with representative error as used in this study. Generally, the

performance of the data assimilation is quite sensitive to the specification of the observation error (Tang et al., 2013). To evaluate the influence of observation error on the optimized emissions, a Sensitivity experiment (EMS7) with doubled observation error was conducted to evaluate the influence of observation error on the optimized emissions. Overall, the spatial distribution of emissions after optimization is was almost the same as that of the EMDA experiment, but with a lowerthe increment is lower (Figure S7), resulting in a weaker estimate of the national total emissions for each species. Thisat is because that the observation error becomes large inflates and, the system will be becomes more convinced certain of the prior emission, and reduces the effect of observation information. Figure 15 shows the time series of simulated and observed daily concentrations and their RMSEs verified against the assimilated sites. The simulations in VEP7 usually performed worse, with larger biases and RMSEs than those of VEP (Figures S8 and S9), especially in most of western and southern China, where posterior emissions are stillwere significantly underestimated. These results usually generally corresponded to sluggish emission changes and large Chi-square statistics (Figure S10), suggesting that too large an observation error that is too large may substantially impact the estimated emissions.

1010

1011

1012

1013

1014

1015

1016

1017

1018

1019

1020

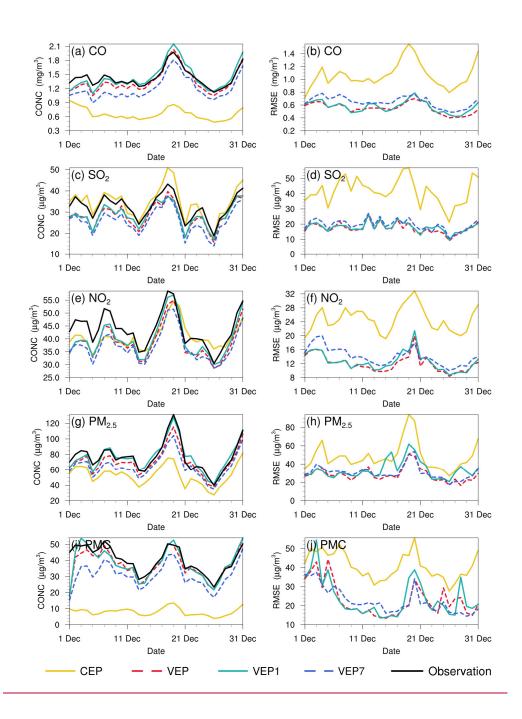
1021

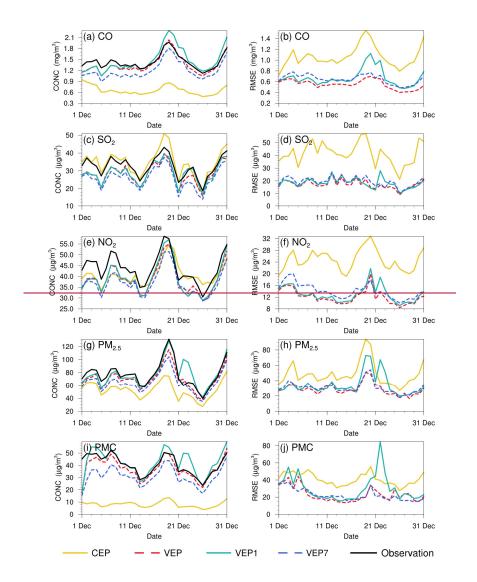
1022

1023

1024

1025





**Figure 15**. Time series of the daily concentrations (CONC, left) and root mean square error (RMSE, right) obtained from CEP, VEP, VEP1, and VEP7. The simulations were verified against the assimilated sites.

## 4.3.4 Impact of the IC optimization of the first window

Many Several studies have shown that there would be indicate large emission discrepancies resulting from the IC errors (Jiang et al., 2013a; Miyazaki et al., 2017; Tang et al., 2013), which means that if the IC is not optimized, the errors of concentrations would be compensated for through the adjustment of emissions. To evaluate the impact of the IC optimization of the first window on the emission inversions, an the EMS8 experiment without the IA step was conducted. Figure 16 shows the time series of the relative differences in the daily posterior emissions of the

five species between the EMDA and EMS8 experiments. It can be found-observed that the IC optimization of IC hads great a significant impact on the emission inversions of long-lived species (i.e., CO). The overall difference in the inverted CO emissions between the two experiments is about was approximately 5.3%, and but can reach 26.1% in the first few windows., the maximum difference can reach 26.1%. For the short-lived species, the IC optimization hads little impact on the emissions; for example, the averaged emission differences of  $SO_2$ ,  $NO_{x_2}$  and PMC in the two experiments are were 0.3%, 0.3%, and 0.9%, respectively. For PPM<sub>2.5</sub>, the average emission difference it is affected not only by the primary emissions, but also by the complex chemistry of its precursors. Therefore, the difference between the two experiments fluctuateds at a certain extent, with overall difference of 2%. It is worth noting that Notably, with the gradual disappearance of the benefit of IC assimilation, the two experiments can reached a unified state after some several windows. For CO, the impact of IA on emission inversion lasteds approximately about half a month. These results indicate that removing the bias of the IC of the first DA window is essential for the subsequent inverse analysis (Jiang et al., 2017).

1040

1041

1042

1043

1044

1045

1046

1047

1048

1049

1050

1051

1052

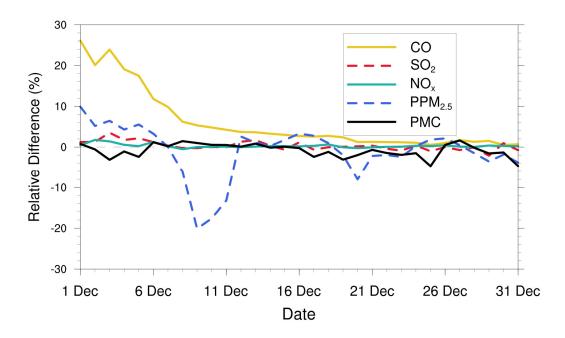
1053

1054

1055

1056

1057



**Figure 16**. Relative differences in CO, SO<sub>2</sub>, NO<sub>2</sub>NO<sub>3</sub>, PPM<sub>2.5</sub>, and PMC emissions (%, the ratio of absolute difference to EMDA) between the EMDA and EMS8.

### 4.3.5 The aAdvantages of the "two-step" scheme

1059

1060 Adjusting the ICs and emissions simultaneously (i.e., "one-step" scheme) has been 1061 applied to constrain prior emissions in many previous studies (e.g., Evensen, 2009; 1062 Kong et al., 2019a). To investigate the impact of different methods on the optimized emissions, a sensitivity test (EMS1) was performed, in which the ICs of each DA 1063 1064 window were <u>also</u> optimized using the <u>3DVAR EnSRF</u> algorithm (Peng et al., 2018; 1065 Schwartz et al., 2014) directly. The spatial localization radius for updating ICs was set 1066 to 90 km in horizontal and 5 layers in vertical closet to the surface with better vertical 1067 mixings. The selections of the horizontal and vertical scales were similar to Kong et al. 1068 (2021) and Tang et al. (2016). We evaluated the optimized ICs of each step, and the 1069 results showed that IC assimilation with EnSRF had good performance (Figure S11). 1070 Compared with our "two-step" method (EMDA), the posterior emissions of EMS1 are were increased by 7.9%, 9.6%, 2.7%, 27.1%, and 22.8%, 7%, 1.4%, 0.6%, 22.2%, and 1071 1072 17.2% higher for CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC, respectively (Figure S12). The 1073 higher emission increase was mainly distributed in the northern China (Figure S13). Overall, there is no significant difference between the two methods for NOx and 1074 SO2, but for CO, it can be clearly seen that the difference increases with the inversion 1075 (Figure S11). We also evaluated the posterior emissions of EMS1 (VEP1) using the 1076 1077 method described in Section 4.1.3. Overall, compared with EMDA, the performance of 1078 EMS1 was worse, with RMSEs of CO, SO<sub>2</sub>, NO<sub>2</sub>, PM<sub>2.5</sub> and PMC increasing from 0.56 mg m<sup>-3</sup> and 17.7, 12.3, 29.6, and 24.6  $\mu$ g m<sup>-3</sup> to 0.58 mg m<sup>-3</sup> and 18.3, 12.9, 34.9, and 1079 25.9 μg m<sup>-3</sup>, respectively (Figure 15). From the perspective of spatial distribution, the 1080 1081 evaluation results become worse in areas where emissions increase (Figure S13). 1082 Additionally, it can be seen from the Figure 15 that the results of the VEP and VEP1 1083 were relatively close at the beginning. However, in the heavy pollution (16–21 1084 December) and later period, the VEP1 with "one-step" inversion emissions had higher concentrations than the observations and larger RMSE than VEP. The results verified 1085 against the independent sites showed a similar situation (Figure S8). This may be 1086 1087 because during the period of heavy pollution, the WRF/CMAQ (offline model) did not consider the feedback process of meteorology and chemistry, resulting in low simulated values. Therefore, the system compensates for the underestimated concentrations caused by the model error through more emissions, resulting in an overestimation of emissions. The accumulation of emission errors in each independent window further leads to the overestimation of concentration after the end of high pollution, especially for species with a long lifetime (e.g. CO). In contrast, using the "two-step" inversion scheme, this overestimation will be corrected quickly in the subsequent inversion to ensure the stability of the system. As mentioned previously, in the "two-step" scheme, the unresolved posterior emission error will beis fed back to the IC of the next window through a sufficient mixed simulation within one day for timely optimization. Meanwhile, the system always maintains the mass balance of the pollutants. In this way Thus, the system updates emissions more consistently and stably. If the emission in one window is overestimated, in this way, it could can be compensated for in the next window with lower estimates. In contrast, when the ICs are optimized assimilating with observations simultaneously at each window, the overestimation will not be corrected and will accumulate to the end-(Figure S14). In addition, the assimilation for initial fields cannot be perfect (Figure S11). As shown in Figure S14, during the heavy pollution episode, there were negative biases in the optimized ICs every day, which lead to a larger positive and a smaller negative emission increment at a certain extent, and result in a larger emission in the end. To remove the effect of this imperfect initial field, we conducted another OSSE experiment (OSSE TRUEIC) using "one-step" scheme, in which the IC of each window was directly taken from the "true" simulation. We further compared the emission error reductions between the OSSE experiment (Section 3) and the OSSE TRUEIC experiment. The results showed that during the last ten days, the error reductions of OSSE TRUEIC were 70.7%, 78.6%, 73.3%, 72.4%, and 63.6% for CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC, respectively, which were smaller than those in the OSSE

1088

1089

1090

1091

1092

1093

1094

1095

1096

1097

1098

1099

1100

1101

1102

1103

1104

1105

1106

1107

1108

1109

1110

1111

1112

1113

1114

1115

window, the inversion performance of "one-step" scheme was still not as good as that of the "two-step" method. We also evaluate the posterior emissions of EMS1 using the same method as shown in Sect. 4.1.3. Overall, compared to the base experiment (EMDA), the performance of EMS1 is significantly worse, with RMSEs of CO, SO<sub>2</sub>, NO<sub>2</sub>, PM<sub>2.5</sub> and PMC increasing from 0.56 mg m<sup>-3</sup>, 17.7, 12.3, 29.6, and 24.6 μg m<sup>-3</sup> to 0.69 mg m<sup>-3</sup>, 18.8, 13.3, 36.8, and 33.3 μg m<sup>-3</sup>, respectively (Figure 15). Additionally, it can be seen from the figure that the results of the two experiments are relatively close at the beginning and during the heavy pollution period (16-21 December). However, after that, the simulated results with "one-step" inversion emissions are significantly higher than the observations, and these large biases continue until the end. The results verified against the independent sites also show a similar situation (Figure S8). The reason may be that during the period of heavy pollution, the WRF-CMAQ (off-line model) does not consider the feedback process of meteorology and chemistry, resulting in low simulations. Therefore, the system will compensate for the underestimated concentrations caused by the model error through more emissions, resulting in the overestimation of emissions. The accumulation of emission error in each independent window further leads to the overestimation of concentration after the end of high pollution, especially for species with a long lifetime (e.g., CO). On the contrary, this overestimation will be corrected quickly in the subsequent inversion using the "twostep" inversion scheme in this study (Figure S11), so as to ensure the stability of the system. Additionally, as shown in section 4.3.1, with the "two-step" scheme, the differences of emissions inverted using MEIC 2012 and 2016 as a priori were only 2.5%, 4.5%, 4.5%, -8.9%, and 3.0% for CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub> and PMC, respectively in the last ten days. We further tested the convergence of the posterior emissions in the "one-step" inversion. the other "one-step" experiment, taking MEIC 2012 as prior emissions, was conducted. However, Except for PPM<sub>2.5</sub>, the relative differences of other species(Figure S12) in posterior emissions between this experiment and the EMS1-did not converge likewere slightly larger than that between EMDA and EMS2 with the "two-step" scheme (Figure

1117

1118

1119

1120

1121

1122

1123

1124

1125

1126

1127

1128

1129

1130

1131

1132

1133

1134

1135

1136

1137

1138

1139

1140

1141

1142

1143

1144

13S16), which further demonstrates underscores the advantages of the "two-step" scheme. It should be noted that the model performance depends on many factors but does not affect the advantage of the "two-step" scheme in emission inversion.

#### 4.4 Discussion

Optimal state estimation using an EnKF relies on the assumption of an unbiased Gaussian —prior error, which is not guaranteed in such highly nonlinear and large biases systems. In this study, some pollutants (e.g., CO, PMC) have very large simulated biases; thus, if a small uncertainty is adopted, the emission bias cannot been fully reduced. while Iif a very large uncertainty is adopted, then the degree of freedom of adjustment is too large, and the inverted daily emissions will fluctuate abnormally. Therefore, we only set a larger prior uncertainty in the first three windows, adoptinged a moderate uncertainty in the following windows, and used a "two-step" inversion scheme and cyclic iteration to gradually correct the emission errors. Figure 10(a) shows the time series of the relative differences between the prior and posterior emissions in each window. There are thewere relatively large adjustments for the emissions in the first three windows, especially for PMC, but the adjustment ranges of the five species after the first 3-three windows are basicallywere within the uncertainty range (e.g., ± 25%), indicating that with this scheme, the EnKF method used in this system still hashad a good performance in emission inversion. The mModel-data mismatch errors not only comesare –from both the emissions, but

The mModel-data mismatch errors not only comesare—from both the emissions, but also from and the inherent model errors arising from the model structure, discretization, parameterizations, and the biases in the simulated meteorological fields. Neglecting model errors would attribute all uncertainties to emissions; and lead to considerable biases in the estimated emissions. In the version of the CMAQ model used in this study, there is are no heterogeneous reactions (Quan et al., 2015; Wang et al., 2017), the parameterization scheme for the formation of secondary organic aerosols (SOA) is imperfect (Carlton et al., 2008; Jiang et al., 2012; Yang et al., 2019), no feedback between chemistry and meteorology is was considered, and we used an idea profile for

chemical lateral boundary conditions. All of the above problems can lead to underestimated concentrations of pollutants, which in turn require more emissions to compensate, leading to overestimations of emissions. In addition, previous studies have shownshowed that the emission of ammonia emissions in the MEIC inventory was are underestimated (Kong et al., 2019b; Paulot et al., 2014; Zhang et al., 2018). Due Owing to lack of ammonia observations, our system does not include emission estimates of ammonia, which means that the concentration of ammonium aerosol was underestimated in this system, also resulting in an overestimation in of the PPM2.5 emission. Wind-blown dust was also not simulated here; thus, the PMC emission inverted in this system do not only come from anthropogenic activities, but also from and natural sources. Although some of these shortcomings could can be solved in the future by updating the CTM model, there will still be errors in each parameterization and each process. Generally In general, a parameter estimation method was used to reduce the model errors, in which, some uncertain parameters were included in the augmented state vector and were optimized synchronously based on the available observations (Brandhorst et al., 2017; Evensen, 2009). However, it is still quite difficult to identify the key uncertain parameters of different species in different models, which generally comes not only from the complex atmospheric chemical model, but also from hundreds of model inputs (Tang et al., 2013). Another method is bias correction, which treats the model error as a bias term, and includes it in the an augmented state vector (Brandhorst et al., 2017; De Lannoy et al., 2007; Keppenne et al., 2005). In addition, the weak-constraint 4D-VarVAR method can also be used to reduce the model errors, which adds a correction term in the model integration to account for the different sources of model error (Sasaki, 1970). Although the reliable diagnosis of model error is stillremains a challenge at present (Laloyaux et al., 2020), it should be considered in an assimilation system. In the future, w we will consider model errors in our system in the future to obtain better emission estimates.

1174

1175

1176

1177

1178

1179

1180

1181

1182

1183

1184

1185

1186

1187

1188

1189

1190

1191

1192

1193

1194

1195

1196

1197

1198

1199

1200

1201

1202

Independent variable localization was adopted to avoid potential spurious correlations across different species in this study. However, the transmission scales for different

species in different regions are still differentdiffer, and a more accurate localization range could can be obtained through backward trajectory analysis. In addition, Although Hamer et al. (2015) successfully used O<sub>2</sub> observations to estimate NOx and VOC emissions within the 4D-var framework within an idealised model, O<sub>3</sub> observations are were not assimilated to improve  $NO_x$  and VOC emissions using crossspecies information. due to the strong nonlinear effects within the O<sub>3</sub>-NO<sub>x</sub>-VOC relationship (Wang et al., 2019b), in which the O<sub>3</sub> concentration and NO<sub>x</sub> (VOC) emissions are were positively correlated in the NO<sub>x</sub> (VOC)-limited region and negatively correlated in the VOC (NO<sub>x</sub>)-limited region (Tang et al., 2011; Wang et al., 2019b). Hamer et al. (2015) successfully used O<sub>3</sub> observations to estimate NO<sub>x</sub> and VOC emissions within the 4DVAR framework within an ideal model. However, the NO<sub>x</sub> emissions are often point or line sources, which are all small compared to the model resolution. With a coarse spatial resolution, the model cannot accurately simulate the relationships between O<sub>3</sub> and its precursors. When assimilating O<sub>3</sub> observations to infer NO<sub>x</sub> or VOC emissions, the inaccurate relationships simulated by model would worsen the inversion of  $NO_x$  emissions (Inness et al., 2015). In general, improving the model resolution can improve the detailed simulation and provide better prior information on O<sub>3</sub>-NO<sub>x</sub>-VOC, but it is still difficult to determine whether the condition is NO<sub>x</sub>-limited or VOC-limited in the real atmosphere using prior emissions (Liu and Shi, 2021). Elbern et al. (2007) emphasized that assimilating O<sub>3</sub> to correct NO<sub>x</sub> or VOC emissions must follow the EKMA framework derived based on observations, otherwise, even if the resolution is improved to sufficiently solve point and line sources, precursor emissions may be still adjusted in an opposite direction. In this study, the spatial resolutions of the prior emission inventory (i.e., MEIC) is  $0.25^{\circ} \times 0.25^{\circ}$ , which is appropriate for modeling at regional scales (Zheng et al., 2017). With this emission inventory, it is unable to accurately simulate the O<sub>3</sub>-NO<sub>x</sub>-VOC relationships. Therefore, to avoid the impact of inaccurate O<sub>3</sub>-NO<sub>x</sub> relationship on emission inversion, in our system, we did not assimilate O<sub>3</sub>, but directly assimilate NO<sub>2</sub> to optimize the NO<sub>x</sub> emissions. This work will be followed up by an ongoing work study using the available

1203

1204

1205

1206

1207

1208

1209

1210

1211

1212

1213

1214

1215

1216

1217

1218

1219

1220

1221

1222

1223

1224

1225

1226

1227

1228

1229

1230

VOC observations.

1232

1233

1234

1235

1236

1237

1238

1239

1240

1241

1242

1243

1244

1245

1246

1247

1248

1249

1250

1251

1252

1253

1254

1255

1256

1257

1258

1259

1260

The initial field. Assuming that NO<sub>2</sub> is underestimated, the NO<sub>2</sub> concentration increases after assimilation, but the VOC concentration remains unchanged, then in the NOx (VOC)-limited region, the subsequent generation of O<sub>3</sub> will increase (decrease); Conversely, the ozone concentration errors caused by assimilating NO2 will also affect the subsequent NOx emission inversion. Similarly, the model may not be able to resolve local scale NO2 well because of uniform distribution of concentration over the whole grid. Therefore, the model is shifted towards a NO<sub>\*</sub> (VOC)-limited regime in high (low) pollution regions, which negatively impacts results by perturbing ozone chemistry in unrealistic ways (Inness et al., 2015). Although we do not assimilate O<sub>3</sub> observation, model resolution still has some influence on inversion results. In our previous study (Feng et al., 2022), we have inferred the NO<sub>x</sub> emissions over YRD in China using NO<sub>2</sub> observations, which has a spatial resolution of 12 km. The study period, assimilated observations, and inversion settings are the same as this study. We compared the posterior emissions of YRD between this study and Feng et al. (2022). The results showed that there was similar spatial distribution of posterior emissions inferred using the two resolutions (36 km vs 12 km) (Figure R17), but the total NO<sub>x</sub> emission in YRD inferred using 36 km resolution was about 8.8% higher than that inferred using 12 km resolution. The differences are mainly caused by meteorological differences at different resolutions. This indicates that coarse model resolution may lead to some overestimation of the inverted emissions. To evaluate the influence of O<sub>3</sub>-NO<sub>x</sub>-VOC relationship change and model resolution on inversion, we also further conducted a nested emission inversion on a densely observed area (the Yangtze River Delta, China) with a grid spacing of 12 km (Feng et al., 2022). The study period is the same as this study. Results showed that the NOx emissions in the Yangtze River Delta retrieved at two resolutions are almost the same (14.7 kt/day vs. 13.4 kt/day), with a difference of 8.8%, indicating that the emissions can be adjusted effectively by RAPAS. In addition, aAs shown previously, the concentrations after DA are obviously were evidently underestimated in western China, indicating that the inverted emissions over these

1261 regions still have large uncertainties because of the sparsity of observations, which that 1262 are spatially insufficient for sampling the inhomogeneity of emissions. Therefore, 1263 further investigations with the joint assimilation of multisource observations (e.g., 1264 satellite) are also underway. 1265 When comparing the performances of the "two-step" and "one-step" schemes, for the 1266 "one-step" scheme, we use a combination assimilation method, namely 3DVAR for the optimizations of initial fields and EnKF for emission inversions in each DA window, 1267 1268 which is similar as Jiang et al., (2017), but different from most previous studies 1269 (Miyazaki et al., 2017; Tang et al., 2013). Because most previous "one-step" 1270 assimilation studies used only one method (i.e., EnKF). This combination method may 1271 cause the comparison less than perfect. However, it should be noted that, even using 1272 the same method (such as EnKF) to optimize the emission of the current window and the initial field of the next window simultaneously (Peng et al., 2018), the initial field 1273 estimation errors will still be mixed in the simulated concentration field, resulting in 1274 1275 unreasonable emission compensation in the next window. In "one-step" scheme, the 1276 essence is to build a good initial field in the high levels. Schwartz et al. (2014) compared 1277 the performances of EnKF and 3DVAR in optimizing initial fields, and found that 1278 3DVAR method can obtain a better initial field than EnKF method. Therefore, we 1279 believe that in this comparison, a combinatorial assimilation approach used in the "onestep" scheme is an acceptable approach, and the conclusion is credible, that the "two-1280 step" scheme has better performances than the "one-step" scheme in emission estimates. 1281 NO<sub>x</sub>NO<sub>x</sub> is mainly emitted by transportation (Li et al., 2017), which can better reflect 1282 1283 the level of economic activities activity to a certain extent. Weekly emission changes 1284 were also explored to verify the performance of the system in depicting emission 1285 changes (Figure S183). Although the "weekend effect" of emissions in China is not 1286 significant (Wang et al., 2014; Wang et al., 2015), the posterior NO<sub>x</sub>NO<sub>x</sub> emission changes showed are ina good agreement with the observations. In our previous studies 1287 1288 (Feng et al., 2020a; Feng et al., 2020b), thise system was successfully applied to optimize NO<sub>x</sub>NO<sub>x</sub> and CO emissions, respectively. The inverted emission changes 1289

were also in line with the time points of epidemic control time points. Additionally, the emission changes can well-reflect the emission migration from developed regions or urban areas to developing regions or surrounding areas over in recent years, which were is consistent with the emission control strategies in China. Although the system does did not consider the model error, resulting in a certain difference between the posterior emission and the actual emissions, the spatiotemporal changes in posterior emissions are were relatively reasonable, which and can be used to monitor emission changes and make inform emission regulations.

## 5 Summary and conclusions

In this study, we developed a Regional multi-Air Pollutant Assimilation System (RAPASv1.0) based on the WRF/CMAQ model, 3DVAR\_algorithm, and EnKF algorithm. RAPAS can quantitatively optimize gridded emissions of CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC on a regional scale by simultaneously assimilating hourly *in\_-situ* measurements of CO, SO<sub>2</sub>, NO<sub>2</sub>, PM<sub>2.5</sub>, and PM<sub>10</sub>. This system includes two subsystems; namely the IA subsystem and the EI subsystem, which optimizes the chemical ICs, and infers the anthropogenic emissions, respectively.

Taking the 2016 Multi-resolution Emission Inventory for China (MEIC 2016) in December as a priori, the emissions of CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC in December 2016 were inferred through by assimilating the corresponding nationwide observations over China. The optimized ICs and posterior emissions were examined against the assimilated and independent observations through parallel forward simulation experiments with and without DA. Sensitivity tests are were also performed to investigate the impact of different inversion processes, prior emissions, prior uncertainties, and observation errors on the emission estimates.

The results show that RAPAS showedhas a good performance in assimilating ground surface *in situ* in-situ observations, with the calculated emission uncertainties reduced by 44.4%, 45.0%, 34.3%, 51.8%, and 56.1% for CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC,

respectively. It can also significantly improve the simulations; the RMSEs of the simulated concentrations with posterior emissions decreased by 40.1—56.3%, and the CORRs increased from 0.26-0.66 to 0.69-0.87 for different species. The OSSE experiment shows showed that the errors of posterior CO, SO<sub>2</sub>, NO<sub>x</sub>, PPM<sub>2.5</sub>, and PMC could be reduced by 78.4%, 86.1%, 78.8%, 77.6%, and 72.0%, respectively. Overall, compared with the prior emissions (MEIC 2016), the posterior emissions increased by 129%, 20%, 5%, and 95% for CO, SO<sub>2</sub>, NO<sub>x2</sub> and PPM<sub>2.5</sub>, respectively. The posterior PMC emissions, which included anthropogenic and natural dust contributions, increased by 1045%. SThe sensitivity tests with different inversion processes with different inversion processes showrevealed that the "two-step" scheme in emission inversion-outperformsed the joint adjustment of ICs and emissions ("one-step" scheme) in emission inversion, especially after heavy pollution. The sSensitivity tests with different prior inventories showed that the observations in China is were sufficient in to infer<del>ring the emissions, and that our system wasis</del> less dependent on prior inventories. Additionally, the sensitivity tests with different prior uncertainties indicated that when the posterior emissions wereare larger than the prior emissions, the emissions decreased/increased with the decreases/increases inef uncertainties because of the different convergence rates. These results demonstrate the advantage of the two-step method in emission inversion in that the inversion errors of the last window caneould be transferred to the current window for further optimization and the robustness of the emissions estimated from RAPAS using the nationwide observations over China. It should be noted that the system usually responds slowly to too small a priori uncertaintiesy or too large observation errors, which may result in large errors in the estimated emissions. In summary, the comprehensive evaluation and sensitivity tests revealed that RAPAS could serve as a useful tool for accurately quantifying the spatial and temporal changes of in multi-species emissions at regional scales and near-real time, which will be helpful for the air pollution control in China, and the other regions around the world with dense

1318

1319

1320

1321

1322

1323

1324

1325

1326

1327

1328

1329

1330

1331

1332

1333

1334

1335

1336

1337

1338

1339

1340

1341

1342

1343

1344

1345

1346

ground observation networks.

## Code and data availability 1347 The codes of RAPAS v1.0 are available at https://doi.org/10.5281/zenodo.5566225. 1348 1349 The WRF model code is open-source code and can be obtained from the WRF Model User's Page (https://www2.mmm.ucar.edu/wrf/users, last access: 25 April 2021). The 1350 CMAQ model is available through an open license as well (https://www.epa.gov/cmaq, 1351 1352 last access: 25 April 2021). The observational and emission data used in this study paper are available at https://doi.org/10.5281/zenodo.4718290 (Feng and Jiang, 2021). 1353 1354 **Author contribution** 1355 SF, FJ, ZW and ZJ developed RAPAS v1.0. SF and FJ designed the research. SF 1356 performed model simulations, analyzed data, and prepared the paper with contributions 1357 from all co-authors. FJ supervised the model development project and assisted in 1358 conceptualization and writing. HW, WH, YS, LZ, YZ, CL, and WJ contributed to the 1359 1360 discussion and improvement of the paper. 1361 **Competing interests** 1362 The authors declare that they have no conflict of interest. 1363 1364 Acknowledgements 1365 This work is supported by the National Key R&D Program of China (Grant No. 1366 1367 2016YFA0600204), the National Natural Science Foundation of China (Grant No. 41907378), and the Nanjing University Innovation and Creative Program for Ph.D. 1368 candidate (Grant No. CXCY19-60). We are grateful to the High Performance 1369 Computing Center (HPCC) of Nanjing University for doing the numerical calculations 1370 in this paper on its blade cluster system, and thank the MEIC team for providing the 1371

prior anthropogenic emissions (http://www.meicmodel.org/).

## 1373 **References**

- Appel, K. W., Pouliot, G. A., Simon, H., Sarwar, G., Pye, H. O. T., Napelenok, S. L., Akhtar, F., and
- 1375 Roselle, S. J.: Evaluation of dust and trace metal estimates from the Community Multiscale Air
- Quality (CMAQ) model version 5.0, Geoscientific Model Development, 6, 883-899,
- 1377 10.5194/gmd-6-883-2013, 2013.
- 1378 Alexe, M., Bergamaschi, P., Segers, A., Detmers, R., Butz, A., Hasekamp, O., Guerlet, S., Parker,
- 1379 R., Boesch, H., Frankenberg, C., Scheepmaker, R. A., Dlugokencky, E., Sweeney, C., Wofsy,
- 1380 S. C., and Kort, E. A.: Inverse modelling of CH4 emissions for 2010-2011 using different
- 1381 satellite retrieval products from GOSAT and SCIAMACHY, Atmospheric Chemistry and
- 1382 Physics, 15, 113-133, 2015.
- 1383 Barbu, A. L., Segers, A. J., Schaap, M., Heemink, A. W., and Builtjes, P. J. H.: A multi-component
- data assimilation experiment directed to sulphur dioxide and sulphate over Europe,
- 1385 Atmospheric Environment, 43, 1622-1631, 2009.
- Bocquet, M.: Parameter-field estimation for atmospheric dispersion: application to the Chernobyl
- 1387 <u>accident using 4D-Var, Quarterly Journal of the Royal Meteorological Society, 138, 664-681,</u>
- 1388 <u>2012.</u>
- 1389 Bocquet, M., Elbern, H., Eskes, H., Hirtl, M., Žabkar, R., Carmichael, G. R., Flemming, J., Inness,
- 1390 A., Pagowski, M., Pérez Camaño, J. L., Saide, P. E., San Jose, R., Sofiev, M., Vira, J., Baklanov,
- 1391 A., Carnevale, C., Grell, G., and Seigneur, C.: Data assimilation in atmospheric chemistry
- 1392 <u>models: current status and future prospects for coupled chemistry meteorology models,</u>
- Atmospheric Chemistry and Physics, 15, 5325-5358, 2015.
- 1394 Bocquet, M. and Sakov, P.: Joint state and parameter estimation with an iterative ensemble Kalman
- 1395 smoother, Nonlinear Processes in Geophysics, 20, 803-818, 2013.
- Basu, S., Guerlet, S., Butz, A., Houweling, S., Hasekamp, O., Aben, I., Krummel, P., Steele, P.,
- Langenfelds, R., Torn, M., Biraud, S., Stephens, B., Andrews, A., and Worthy, D.: Global CO2
- 1398 fluxes estimated from GOSAT retrievals of total column CO2, Atmospheric Chemistry and
- 1399 Physics, 13, 8695-8717, 2013.
- 1400 Bauwens, M., Compernolle, S., Stavrakou, T., Müller, J.-F., van Gent, J., Eskes, H., Levelt, P. F.,
- 1401 van der A, R., Veefkind, J. P., Vlietinck, J., Yu, H., and Zehner, C.: Impact of Coronavirus
- Outbreak on NO2 Pollution Assessed Using TROPOMI and OMI Observations, 47,
- 1403 e2020GL087978, 10.1029/2020gl087978, 2020.
- 1404 Bierman: Factorization methods for Discrete Sequential estimation, Academic Press, 1977.
- 1405 Binkowski, F. S. and Roselle, S. J.: Models-3 community multiscale air quality (CMAQ) model
- 1406 aerosol component 1. Model description, Journal of Geophysical Research-Atmospheres, 108,
- 1407 10.1029/2001jd001409, 2003.
- 1408 Brandhorst, N., Erdal, D., and Neuweiler, I.: Soil moisture prediction with the ensemble Kalman
- filter: Handling uncertainty of soil hydraulic parameters, Advances in Water Resources, 110,
- 1410 360-370, 2017.

- 1411 Bruhwiler, L. M. P., Michalak, A. M., Peters, W., Baker, D. F., and Tans, P.: An improved Kalman
- Smoother for atmospheric inversions, Atmos. Chem. Phys., 5, 2691-2702, 10.5194/acp-5-
- 1413 2691-2005, 2005.
- 1414 Carlton, A. G., Turpin, B. J., Altieri, K. E., Seitzinger, S. P., Mathur, R., Roselle, S. J., and Weber,
- 1415 R. J.: CMAQ Model Performance Enhanced When In-Cloud Secondary Organic Aerosol is
- 1416 Included: Comparisons of Organic Carbon Predictions with Measurements, Environmental
- 1417 Science & Technology, 42, 8798-8802, 2008
- 1418 Chen, D., Liu, Z., Ban, J., and Chen, M.: The 2015 and 2016 wintertime air pollution in China: SO2
- emission changes derived from a WRF-Chem/EnKF coupled data assimilation system,
- 1420 Atmospheric Chemistry and Physics, 19, 8619-8650, 10.5194/acp-19-8619-2019, 2019.
- 1421 Chen, D., Liu, Z., Fast, J., and Ban, J.: Simulations of sulfate-nitrate-ammonium (SNA) aerosols
- during the extreme haze events over northern China in October 2014, Atmospheric Chemistry
- and Physics, 16, 10707-10724, 10.5194/acp-16-10707-2016, 2016.
- 1424 Chevallier, F., Bréon, F.-M., and Rayner, P. J.: Contribution of the Orbiting Carbon Observatory to
- the estimation of CO2 sources and sinks: Theoretical study in a variational data assimilation
- 1426 framework, 112, 10.1029/2006JD007375, 2007.
- 1427 Clements, A. L., Fraser, M. P., Upadhyay, N., Herckes, P., Sundblom, M., Lantz, J., and Solomon,
- P. A.: Chemical characterization of coarse particulate matter in the Desert Southwest Pinal
- 1429 County Arizona, USA, Atmospheric Pollution Research, 5, 52-61, 10.5094/apr.2014.007, 2014.
- 1430 Clements, N., Hannigan, M. P., Miller, S. L., Peel, J. L., and Milford, J. B.: Comparisons of urban
- and rural PM10-2.5 and PM2.5 mass concentrations and semi-volatile fractions in northeastern
- 1432 Colorado, Atmospheric Chemistry and Physics, 16, 7469-7484, 10.5194/acp-16-7469-2016,
- 1433 2016.
- 1434 Daley, R.: Atmospheric Data Assimilation (gtSpecial IssueltData Assimilation in Meteology and
- Oceanography: Theory and Practice), Journal of the Meteorological Society of Japan. Ser. II,
- 1436 75, 319-329, 1997.
- 1437 Derber, J. C.: A VARIATIONAL CONTINUOUS ASSIMILATION TECHNIQUE, Monthly
- 1438 Weather Review, 117, 2437-2446, 1989.
- de Foy, B., Lu, Z., Streets, D. G., Lamsal, L. N., and Duncan, B. N.: Estimates of power plant NOx
- emissions and lifetimes from OMI NO2 satellite retrievals, Atmospheric Environment, 116, 1-
- 1441 11, 10.1016/j.atmosenv.2015.05.056, 2015.
- 1442 De Lannoy, G. J. M., Houser, P. R., Pauwels, V. R. N., and Verhoest, N. E. C.: State and bias
- estimation for soil moisture profiles by an ensemble Kalman filter: Effect of assimilation depth
- 1444 and frequency, 43, 2007.
- Ding, J., van der A, R. J., Mijling, B., Levelt, P. F., and Hao, N.: NOx emission estimates during the
- 2014 Youth Olympic Games in Nanjing, Atmospheric Chemistry and Physics, 15, 9399-9412,
- 1447 10.5194/acp-15-9399-2015, 2015.
- 1448 Elbern, H., Strunk, A., Schmidt, H., and Talagrand, O.: Emission rate and chemical state estimation

- by 4-dimensional variational inversion, Atmospheric Chemistry and Physics, 7, 3749-3769,
- 1450 10.5194/acp-7-3749-2007, 2007.
- 1451 Evensen, G.: The Ensemble Kalman Filter for Combined State and Parameter Estimation MONTE
- 1452 CARLO TECHNIQUES FOR DATA ASSIMILATION IN LARGE SYSTEMS, Ieee Control
- 1453 Systems Magazine, 29, 83-104, 10.1109/mcs.2009.932223, 2009.
- 1454 Feng, S., Jiang, F., Jiang, Z., Wang, H., Cai, Z., and Zhang, L.: Impact of 3DVAR assimilation of
- surface PM2.5 observations on PM2.5 forecasts over China during wintertime, Atmospheric
- Environment, 187, 34-49, 10.1016/j.atmosenv.2018.05.049, 2018.
- 1457 Feng, S., Jiang, F., Wang, H., Shen, Y., Zheng, Y., Zhang, L., Lou, C., and Ju, W.: Anthropogenic
- emissions estimated using surface observations and their impacts on PM2.5 source
- apportionment over the Yangtze River Delta, China, Science of The Total Environment, 828,
- 1460 154522, 2022
- 1461 Feng, S., Jiang, F., Wu, Z., Wang, H., Ju, W., and Wang, H.: CO Emissions Inferred From Surface
- 1462 CO Observations Over China in December 2013 and 2017, Journal of Geophysical Research-
- 1463 Atmospheres, 125, 10.1029/2019jd031808, 2020a.
- 1464 Feng, S., Jiang, F., Wang, H., Wang, H., Ju, W., Shen, Y., Zheng, Y., Wu, Z., and Ding, A.: NOx
- Emission Changes Over China During the COVID-19 Epidemic Inferred From Surface NO2
- Observations, Geophysical Research Letters, 47, 10.1029/2020gl090080, 2020b.
- 1467 Feng, S. and Jiang, F.: Anthropogenic air pollutant emissions over China inferred by Regional multi-
- 1468 Air Pollutant Assimilation System (RAPAS v1.0), Zenodo, 10.5281/zenodo.4718290, 2021.
- Gaspari, G. and Cohn, S. E.: Construction of correlation functions in two and three dimensions,
- Quarterly Journal of the Royal Meteorological Society, 125, 723-757, 10.1256/smsqj.55416,
- 1471 1999.
- Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L. K., and Wang,
- 1473 X.: The Model of Emissions of Gases and Aerosols from Nature version 2.1 (MEGAN2.1): an
- extended and updated framework for modeling biogenic emissions, Geoscientific Model
- 1475 Development, 5, 1471-1492, 10.5194/gmd-5-1471-2012, 2012.
- 1476 Gurney, K. R., Law, R. M., Denning, A. S., Rayner, P. J., Pak, B. C., Baker, D., Bousquet, P.,
- Bruhwiler, L., Chen, Y. H., Ciais, P., Fung, I. Y., Heimann, M., John, J., Maki, T., Maksyutov,
- 1478 S., Peylin, P., Prather, M., and Taguchi, S.: Transcom 3 inversion intercomparison: Model mean
- results for the estimation of seasonal carbon sources and sinks, Global Biogeochemical Cycles,
- 1480 18, 10.1029/2003gb002111, 2004.
- 1481 He, W., van der Velde, I. R., Andrews, A. E., Sweeney, C., Miller, J., Tans, P., van der Laan-Luijkx,
- 1482 I. T., Nehrkorn, T., Mountain, M., Ju, W., Peters, W., and Chen, H.: CTDAS-Lagrange v1.0: a
- high-resolution data assimilation system for regional carbon dioxide observations,
- 1484 Geoscientific Model Development, 11, 3515-3536, 10.5194/gmd-11-3515-2018, 2018.
- 1485 Hinds, W.C.: Aerosol Technology: Properties, Behavior, and Measurement of Airborne Particles.
- 1486 New York: John Wiley, 1982.

- 1487 Houtekamer, P. L. and Mitchell, H. L.: A sequential ensemble Kalman filter for atmospheric data
- 1488 assimilation, Monthly Weather Review, 129, 123-137, 10.1175/1520-
- 1489 0493(2001)129<0123:asekff>2.0.co;2, 2001.
- 1490 Houtekamer, P. L. and Zhang, F.: Review of the Ensemble Kalman Filter for Atmospheric Data
- 1491 Assimilation, Monthly Weather Review, 144, 4489-4532, 10.1175/mwr-d-15-0440.1, 2016.
- 1492 Inness, A., Blechschmidt, A. M., Bouarar, I., Chabrillat, S., Crepulja, M., Engelen, R. J., Eskes, H.,
- Flemming, J., Gaudel, A., Hendrick, F., Huijnen, V., Jones, L., Kapsomenakis, J., Katragkou,
- 1494 E., Keppens, A., Langerock, B., de Maziere, M., Melas, D., Parrington, M., Peuch, V. H.,
- Razinger, M., Richter, A., Schultz, M. G., Suttie, M., Thouret, V., Vrekoussis, M., Wagner, A.,
- 1496 and Zerefos, C.: Data assimilation of satellite-retrieved ozone, carbon monoxide and nitrogen
- dioxide with ECMWF's Composition-IFS, Atmospheric Chemistry and Physics, 15, 5275-5303,
- 1498 2015.
- Jiang, F., Liu, Q., Huang, X., Wang, T., Zhuang, B., and Xie, M.: Regional modeling of secondary
- 1500 organic aerosol over China using WRF/Chem, Journal of Aerosol Science, 43, 57-73,
- 1501 10.1016/j.jaerosci.2011.09.003, 2012a.
- 1502 Jiang, F., Zhou, P., Liu, Q., Wang, T., Zhuang, B., and Wang, X.: Modeling tropospheric ozone
- formation over East China in springtime, Journal of Atmospheric Chemistry, 69, 303-319,
- 1504 10.1007/s10874-012-9244-3, 2012b.
- 1505 Jiang, F., Wang, H. M., Chen, J. M., Machida, T., Zhou, L. X., Ju, W. M., Matsueda, H., and Sawa,
- 1506 Y.: Carbon balance of China constrained by CONTRAIL aircraft CO2 measurements,
- 1507 Atmospheric Chemistry and Physics, 14, 10133-10144, 10.5194/acp-14-10133-2014, 2014.
- 1508 Jiang, F., Wang, H., Chen, J. M., Ju, W., Tian, X., Feng, S., Li, G., Chen, Z., Zhang, S., Lu, X., Liu,
- 1509 J., Wang, H., Wang, J., He, W., and Wu, M.: Regional CO2 fluxes from 2010 to 2015 inferred
- from GOSAT XCO2 retrievals using a new version of the Global Carbon Assimilation System,
- 1511 Atmos. Chem. Phys., 21, 1963-1985, 10.5194/acp-21-1963-2021, 2021.
- 1512 Jiang, W., Smyth, S., Giroux, E., Roth, H., and Yin, D.: Differences between CMAQ fine mode
- 1513 particle and PM2.5 concentrations and their impact on model performance evaluation in the
- 1514 lower Fraser valley, Atmospheric Environment, 40, 4973-4985,
- 1515 10.1016/j.atmosenv.2005.10.069, 2006.
- 1516 Jiang, Z., Jones, D. B. A., Worden, H. M., Deeter, M. N., Henze, D. K., Worden, J., Bowman, K. W.,
- Brenninkmeijer, C. A. M., and Schuck, T. J.: Impact of model errors in convective transport on
- 1518 CO source estimates inferred from MOPITT CO retrievals, Journal Of Geophysical Research-
- 1519 Atmospheres, 118, 2073-2083, 2013a.
- 1520 Jiang, Z., Liu, Z., Wang, T., Schwartz, C. S., Lin, H.-C., and Jiang, F.: Probing into the impact of
- 1521 3DVAR assimilation of surface PM10 observations over China using process analysis, Journal
- of Geophysical Research: Atmospheres, 118, 6738-6749, 10.1002/jgrd.50495, 2013b.
- 1523 Jiang, Z., Worden, J. R., Worden, H., Deeter, M., Jones, D. B. A., Arellano, A. F., and Henze, D. K.:
- A 15-year record of CO emissions constrained by MOPITT CO observations, Atmospheric
- 1525 Chemistry And Physics, 17, 4565-4583, 10.5194/acp-17-4565-2017, 2017.

- 1526 Jin, J., Lin, H. X., Heemink, A., and Segers, A.: Spatially varying parameter estimation for dust
- emissions using reduced-tangent-linearization 4DVar, Atmospheric Environment, 187, 358-
- 1528 373, 10.1016/j.atmosenv.2018.05.060, 2018.
- Kahnert, M.: Variational data analysis of aerosol species in a regional CTM: background error
- 1530 covariance constraint and aerosol optical observation operators, Tellus B, 60, 2008.
- Kang, J.-S., Kalnay, E., Miyoshi, T., Liu, J., and Fung, I.: Estimation of surface carbon fluxes with
- an advanced data assimilation methodology, 117, 10.1029/2012JD018259, 2012.
- 1533 Keppenne, C. L., Rienecker, M. M., Kurkowski, N. P., and Adamec, D. A.: Ensemble Kalman filter
- assimilation of temperature and altimeter data with bias correction and application to seasonal
- prediction, Nonlin. Processes Geophys., 12, 491-503, 2005.
- 1536 Kleist, D. T., Parrish, D. F., Derber, J. C., Treadon, R., Wu, W.-S., and Lord, S.: Introduction of the
- 1537 GSI into the NCEP Global Data Assimilation System, Weather and Forecasting, 24, 1691-1705,
- 1538 10.1175/2009waf2222201.1, 2009.
- 1539 Kong, L., Tang, X., Zhu, J., Wang, Z., Pan, Y., Wu, H., Wu, L., Wu, Q., He, Y., Tian, S., Xie, Y., Liu,
- 2., Sui, W., Han, L., and Carmichael, G.: Improved Inversion of Monthly Ammonia Emissions
- in China Based on the Chinese Ammonia Monitoring Network and Ensemble Kalman Filter,
- 1542 Environmental Science & Technology, 53, 12529-12538, 10.1021/acsest.9b02701, 2019a.
- Kong, L., Tang, X., Zhu, J., Wang, Z., Fu, J. S., Wang, X., Itahashi, S., Yamaji, K., Nagashima, T.,
- 1544 Lee, H. J., Kim, C. H., Lin, C. Y., Chen, L., Zhang, M., Tao, Z., Li, J., Kajino, M., Liao, H.,
- Sudo, K., Wang, Y., Pan, Y., Tang, G., Li, M., Wu, Q., Ge, B., and Carmichael, G. R.: Evaluation
- and uncertainty investigation of the NO2, CO and NH3 modeling over China under the
- 1547 framework of MICS-Asia III, Atmos. Chem. Phys. Discuss., 2019, 1-33, 10.5194/acp-2018-
- 1548 1158, 2019b.
- 1549 Kurokawa, J.-i., Yumimoto, K., Uno, I., and Ohara, T.: Adjoint inverse modeling of NOx emissions
- 1550 over eastern China using satellite observations of NO2 vertical column densities, Atmospheric
- 1551 Environment, 43, 1878-1887, 10.1016/j.atmosenv.2008.12.030, 2009.
- Laloyaux, P., Bonavita, M., Chrust, M., and Gürol, S.: Exploring the potential and limitations of
- 1553 weak-constraint 4D-Var, Quarterly Journal of the Royal Meteorological Society, 146, 4067-
- 1554 4082, 2020
- 1555 Li, J.-d., Deng, Q.-h., Lu, C., and Huang, B.-l.: Chemical compositions and source apportionment
- 1556 of atmospheric PM10 in suburban area of Changsha, China, Journal of Central South
- 1557 University of Technology, 17, 509-515, 2010.
- 1558 Li, M., Zhang, Q., Kurokawa, J.-i., Woo, J.-H., He, K., Lu, Z., Ohara, T., Song, Y., Streets, D. G.,
- 1559 Carmichael, G. R., Cheng, Y., Hong, C., Huo, H., Jiang, X., Kang, S., Liu, F., Su, H., and Zheng,
- B.: MIX: a mosaic Asian anthropogenic emission inventory under the international
- 1561 collaboration framework of the MICS-Asia and HTAP, Atmospheric Chemistry and Physics,
- 1562 17, 935-963, 10.5194/acp-17-935-2017, 2017.
- Liu, C. and Shi, K.: A review on methodology in O3-NOx-VOC sensitivity study, Environmental
- 1564 <u>Pollution, 291, 118249, 2021.</u>

- Liu, Y., Kalnay, E., Zeng, N., Asrar, G., Chen, Z., and Jia, B.: Estimating surface carbon fluxes based
- on a local ensemble transform Kalman filter with a short assimilation window and a long
- 1567 <u>observation window: an observing system simulation experiment test in GEOS-Chem 10.1,</u>
- 1568 <u>Geoscientific Model Development</u>, 12, 2899-2914, 2019.
- Liu, Z., Liu, Q., Lin, H.-C., Schwartz, C. S., Lee, Y.-H., and Wang, T.: Three-dimensional variational
- assimilation of MODIS aerosol optical depth: Implementation and application to a dust storm
- 1571 over East Asia, Journal of Geophysical Research: Atmospheres, 116, n/a-n/a,
- 1572 10.1029/2011jd016159, 2011.
- Lorenc, A. C.: Modelling of error covariances by 4D-Var data assimilation, Quarterly Journal of the
- 1574 Royal Meteorological Society, 129, 3167-3182, 2003.
- Hamer, P. D., Bowman, K. W., Henze, D. K., Attie, J. L., and Marecal, V.: The impact of observing
- characteristics on the ability to predict ozone under varying polluted photochemical regimes,
- 1577 Atmospheric Chemistry and Physics, 15, 10645-10667, 2015.
- 1578 Ma, C., Wang, T., Mizzi, A. P., Anderson, J. L., Zhuang, B., Xie, M., and Wu, R.: Multiconstituent
- Data Assimilation With WRF-Chem/DART: Potential for Adjusting Anthropogenic Emissions
- and Improving Air Quality Forecasts Over Eastern China, 124, 7393-7412,
- 1581 10.1029/2019jd030421, 2019.
- Meirink, J. F., Bergamaschi, P., and Krol, M. C.: Four-dimensional variational data assimilation for
- inverse modelling of atmospheric methane emissions: method and comparison with synthesis
- 1584 <u>inversion, Atmospheric Chemistry and Physics, 8, 6341-6353, 2008.</u>
- Meirink, J. F., Eskes, H. J., and Goede, A. P. H.: Sensitivity analysis of methane emissions derived
- 1586 from SCIAMACHY observations through inverse modelling, Atmospheric Chemistry and
- 1587 Physics, 6, 1275-1292, 10.5194/acp-6-1275-2006, 2006.
- 1588 Maybeck: Stochastic Models, Estimation and Control Academic Press, 1979.
- 1589 Miyazaki, K. and Eskes, H.: Constraints on surface NOx emissions by assimilating satellite
- observations of multiple species, Geophysical Research Letters, 40, 4745-4750,
- 1591 10.1002/grl.50894, 2013.
- 1592 Miyazaki, K., Eskes, H. J., and Sudo, K.: Global NOx emission estimates derived from an
- assimilation of OMI tropospheric NO2 columns, Atmospheric Chemistry and Physics, 12,
- 2263-2288, 10.5194/acp-12-2263-2012, 2012a.
- 1595 Miyazaki, K., Eskes, H. J., Sudo, K., Takigawa, M., van Weele, M., and Boersma, K. F.:
- Simultaneous assimilation of satellite NO2, O-3, CO, and HNO3 data for the analysis of
- tropospheric chemical composition and emissions, Atmospheric Chemistry and Physics, 12,
- 1598 9545-9579, 10.5194/acp-12-9545-2012, 2012b.
- 1599 Miyazaki, K., Eskes, H., Sudo, K., Boersma, K. F., Bowman, K., and Kanaya, Y.: Decadal changes
- in global surface NOx emissions from multi-constituent satellite data assimilation,
- Atmospheric Chemistry and Physics, 17, 807-837, 2017.
- 1602 Mizzi, A. P., Edwards, D. P., and Anderson, J. L.: Assimilating compact phase space retrievals

- 1603 (CPSRs): comparison with independent observations (MOZAIC in situ and IASI retrievals)
  1604 and extension to assimilation of truncated retrieval profiles, Geoscientific Model Development,
  1605 11, 3727-3745, 2018.
- Monteil, G., Houweling, S., Butz, A., Guerlet, S., Schepers, D., Hasekamp, O., Frankenberg, C.,
  Scheepmaker, R., Aben, I., and Rockmann, T.: Comparison of CH4 inversions based on 15
  months of GOSAT and SCIAMACHY observations, Journal of Geophysical ResearchAtmospheres, 118, 11807-11823, 2013.
- Muller, J. F. and Stavrakou, T.: Inversion of CO and NOx emissions using the adjoint of the IMAGES model, Atmospheric Chemistry and Physics, 5, 1157-1186, 2005.
- Nassar, R., Jones, D. B. A., Kulawik, S. S., Worden, J. R., Bowman, K. W., Andres, R. J., Suntharalingam, P., Chen, J. M., Brenninkmeijer, C. A. M., Schuck, T. J., Conway, T. J., and Worthy, D. E.: Inverse modeling of CO2 sources and sinks using satellite observations of CO2 from TES and surface flask measurements, Atmospheric Chemistry and Physics, 11, 6029-6047, 2011.
- Navon, I. M.: Practical and theoretical aspects of adjoint parameter estimation and identifiability in meteorology and oceanography, Dynamics of Atmospheres and Oceans, 27, 55-79, 1998.
- Parrish, D. F. and Derber, J. C.: The National Meteorological Center's spectral statisticalinterpolation analysis system, Monthly Weather Review, 120, 1747-1763, 10.1175/1520-0493(1992)120<1747:tnmcss>2.0.co;2, 1992.
- Paulot, F., Jacob, D. J., Pinder, R. W., Bash, J. O., Travis, K., and Henze, D. K.: Ammonia emissions in the United States, European Union, and China derived by high-resolution inversion of ammonium wet deposition data: Interpretation with a new agricultural emissions inventory (MASAGE\_NH3), Journal of Geophysical Research-Atmospheres, 119, 4343-4364, 2014.
- Peng, Z., Liu, Z., Chen, D., and Ban, J.: Improving PM<sub&gt; 2. 5&lt;/sub&gt; forecast over China by the joint adjustment of initial conditions and source emissions with an ensemble Kalman filter, Atmospheric Chemistry and Physics, 17, 4837-4855, 10.5194/acp-17-4837-2017, 2017.
- Peng, Z., Lei, L., Liu, Z., Su, J., Ding, A., Ban, J., Chen, D., Kou, X., and Chu, K.: The impact of multi-species surface chemical observation assimilation on air quality forecasts in China, Atmospheric Chemistry and Physics, 18, 10.5194/acp-18-17387-2018, 2018.
- Peters, W., Jacobson, A. R., Sweeney, C., Andrews, A. E., Conway, T. J., Masarie, K., Miller, J. B.,
   Bruhwiler, L. M. P., Petron, G., Hirsch, A. I., Worthy, D. E. J., van der Werf, G. R., Randerson,
   J. T., Wennberg, P. O., Krol, M. C., and Tans, P. P.: An atmospheric perspective on North
   American carbon dioxide exchange: CarbonTracker, Proceedings of the National Academy of
   Sciences of the United States of America, 104, 18925-18930, 10.1073/pnas.0708986104, 2007.
- Peylin, P., Rayner, P. J., Bousquet, P., Carouge, C., Hourdin, F., Heinrich, P., Ciais, P., and contributors, A.: Daily CO2 flux estimates over Europe from continuous atmospheric measurements: 1, inverse methodology, Atmospheric Chemistry and Physics, 5, 3173-3186, 10.5194/acp-5-3173-2005, 2005.

- Purser, R. J., Wu, W. S., Parrish, D. F., and Roberts, N. M.: Numerical aspects of the application of
- recursive filters to variational statistical analysis. Part I: Spatially homogeneous and isotropic
- 1644 Gaussian covariances, Monthly Weather Review, 131, 1524-1535, 10.1175//1520-
- 1645 0493(2003)131<1524:naotao>2.0.co;2, 2003.
- Quan, J., Liu, Q., Li, X., Gao, Y., Jia, X., Sheng, J., Liu, Y., 2015. Effect of heterogeneous aqueous
- reactions on the secondary formation of inorganic aerosols during haze events. Atmospheric
- 1648 Environment 122, 306-312.
- Rabier, F., McNally, A., Andersson, E., Courtier, P., Unden, P., Eyre, J., Hollingsworth, A., and
- Bouttier, F.: The ECMWF implementation of three-dimensional variational assimilation (3D-
- Var). II: Structure functions, Quarterly Journal Of the Royal Meteorological Society, 124,
- 1652 1809-1829, 10.1256/smsqj.55002, 1998.
- Reichle, R. H., McLaughlin, D. B., and Entekhabi, D.: Hydrologic data assimilation with the
- ensemble Kalman filter, Monthly Weather Review, 130, 103-114, 2002.
- 1655 Richardson, H., Basu, S., and Holtslag, A. A. M.: Improving Stable Boundary-Layer Height
- 1656 Estimation Using a Stability-Dependent Critical Bulk Richardson Number, Boundary-Layer
- 1657 Meteorology, 148, 93-109, 2013.
- 1658 Ruiz, J. and Pulido, M.: Parameter Estimation Using Ensemble-Based Data Assimilation in the
- Presence of Model Error, Monthly Weather Review, 143, 1568-1582, 2015.
- Sarwar, G., Simon, H., Bhave, P., and Yarwood, G.: Examining the impact of heterogeneous nitryl
- 1661 chloride production on air quality across the United States, Atmospheric Chemistry and
- Physics, 12, 6455-6473, 10.5194/acp-12-6455-2012, 2012.
- 1663 Sasaki, Y.: SOME BASIC FORMALISMS IN NUMERICAL VARIATIONAL ANALYSIS,
- 1664 Monthly Weather Review, 98, 875-&, 1970.
- Schneising, O., Buchwitz, M., Burrows, J. P., Bovensmann, H., Bergamaschi, P., and Peters, W.:
- Three years of greenhouse gas column-averaged dry air mole fractions retrieved from satellite
- Part 2: Methane, Atmospheric Chemistry and Physics, 9, 443-465, 2009.
- 1668 Schwartz, C. S., Liu, Z., Lin, H.-C., and Cetola, J. D.: Assimilating aerosol observations with a
- 1669 "hybrid" variational-ensemble data assimilation system, Journal Of Geophysical Research-
- 1670 Atmospheres, 119, 4043-4069, 10.1002/2013jd020937, 2014.
- 1671 Sekiyama, T. T., Tanaka, T. Y., Shimizu, A., and Miyoshi, T.: Data assimilation of CALIPSO aerosol
- observations, Atmospheric Chemistry and Physics, 10, 39-49, 10.5194/acp-10-39-2010, 2010.
- 1673 Shen, Y., Jiang, F., Feng, S., Zheng, Y., Cai, Z., and Lyu, X.: Impact of weather and emission changes
- on NO2 concentrations in China during 2014–2019, Environmental Pollution, 269, 116163,
- 1675 10.1016/j.envpol.2020.116163, 2021.
- 1676 Shi, X. and Brasseur, G. P.: The Response in Air Quality to the Reduction of Chinese Economic
- 1677 Activities During the COVID-19 Outbreak, 47, e2020GL088070, 10.1029/2020gl088070,
- 1678 2020.
- Stanevich, I., Jones, D. B. A., Strong, K., Keller, M., Henze, D. K., Parker, R. J., Boesch, H., Wunch,

- 1680 D., Notholt, J., Petri, C., Warneke, T., Sussmann, R., Schneider, M., Hase, F., Kivi, R.,
- Deutscher, N. M., Velazco, V. A., Walker, K. A., and Deng, F.: Characterizing model errors in
- chemical transport modeling of methane: using GOSAT XCH4 data with weak-constraint four-
- dimensional variational data assimilation, Atmospheric Chemistry and Physics, 21, 9545-9572,
- 1684 2021.
- Stavrakou, T., Müller, J.-F., Boersma, K. F., De Smedt, I., and van der A, R. J.: Assessing the
- distribution and growth rates of NOx emission sources by inverting a 10-year record of NO2
- satellite columns, 35, 10.1029/2008gl033521, 2008.
- Sun, A. Y., Morris, A., and Mohanty, S.: Comparison of deterministic ensemble Kalman filters for
- assimilating hydrogeological data, Advances in Water Resources, 32, 280-292,
- 1690 10.1016/j.advwatres.2008.11.006, 2009.
- Takagi, H., Saeki, T., Oda, T., Saito, M., Valsala, V., Belikov, D., Saito, R., Yoshida, Y., Morino, I.,
- Uchino, O., Andres, R. J., Yokota, T., and Maksyutov, S.: On the Benefit of GOSAT
- Observations to the Estimation of Regional CO<sub>2</sub> Fluxes, SOLA, 7, 161-164,
- 1694 10.2151/sola.2011-041, 2011.
- 1695 Tang, X., Zhu, J., Wang, Z. F., and Gbaguidi, A.: Improvement of ozone forecast over Beijing based
- on ensemble Kalman filter with simultaneous adjustment of initial conditions and emissions,
- Atmospheric Chemistry And Physics, 11, 12901-12916, 10.5194/acp-11-12901-2011, 2011.
- 1698 Tang, X., Zhu, J., Wang, Z. F., Wang, M., Gbaguidi, A., Li, J., Shao, M., Tang, G. Q., and Ji, D. S.:
- Inversion of CO emissions over Beijing and its surrounding areas with ensemble Kalman filter,
- 1700 Atmospheric Environment, 81, 676-686, 10.1016/j.atmosenv.2013.08.051, 2013.
- 1701 Wang, C., Lei, L., Tan, Z.-M., and Chu, K.: Adaptive Localization for Tropical Cyclones With
- 1702 Satellite Radiances in an Ensemble Kalman Filter, Frontiers in Earth Science, 8,
- 1703 10.3389/feart.2020.00039, 2020.
- Wang, H., Jiang, F., Wang, J., Ju, W., and Chen, J. M.: Terrestrial ecosystem carbon flux estimated
- using GOSAT and OCO-2 XCO2 retrievals, Atmospheric Chemistry and Physics, 19, 12067-
- 1706 12082, 2019a.
- Wang, N., Lyu, X., Deng, X., Huang, X., Jiang, F., and Ding, A.: Aggravating O3 pollution due to
- 1708 NOx emission control in eastern China, Science of The Total Environment, 677, 732-744,
- 1709 2019b.
- 1710 Wang, Y. H., Hu, B., Ji, D. S., Liu, Z. R., Tang, G. Q., Xin, J. Y., Zhang, H. X., Song, T., Wang, L.
- 1711 L., Gao, W. K., Wang, X. K., and Wang, Y. S.: Ozone weekend effects in the Beijing-Tianjin-
- Hebei metropolitan area, China, Atmospheric Chemistry and Physics, 14, 2419-2429, 2014.
- Wang, Z., Li, Y., Dong, X., Sun, R., Sun, N., and Pan, L.: Analysis on weekend effect of air
- pollutants in urban atmosphere of Beijing, Journal of University of Chinese Academy of
- 1715 Sciences, 32, 843-850, 2015.
- 1716 Wang, Z., Wang, W., Tham, Y.J., Li, Q., Wang, H., Wen, L., Wang, X., Wang, T., 2017. Fast
- 1717 heterogeneous N2O5 uptake and ClNO2 production in power plant and industrial plumes
- observed in the nocturnal residual layer over the North China Plain. Atmospheric Chemistry

- 1719 and Physics 17, 12361-12378.
- Wecht, K. J., Jacob, D. J., Sulprizio, M. P., Santoni, G. W., Wofsy, S. C., Parker, R., Boesch, H., and
- Worden, J.: Spatially resolving methane emissions in California: constraints from the CalNex
- aircraft campaign and from present (GOSAT, TES) and future (TROPOMI, geostationary)
- satellite observations, Atmospheric Chemistry and Physics, 14, 8173-8184, 2014.
- Wu, H., Tang, X., Wang, Z., Wu, L., Li, J., Wang, W., Yang, W., and Zhu, J.: High-spatiotemporal-
- 1725 resolution inverse estimation of CO and NOx emission reductions during emission control
- periods with a modified ensemble Kalman filter, Atmospheric Environment, 236,
- 1727 10.1016/j.atmosenv.2020.117631, 2020.
- Wu, W. S., Purser, R. J., and Parrish, D. F.: Three-dimensional variational analysis with spatially
- inhomogeneous covariances, Monthly Weather Review, 130, 2905-2916, 10.1175/1520-
- 1730 0493(2002)130<2905:tdvaws>2.0.co;2, 2002.
- 1731 Yang, W., Li, J., Wang, W., Li, J., Ge, M., Sun, Y., Chen, X., Ge, B., Tong, S., Wang, Q., and Wang,
- 2.1 Investigating secondary organic aerosol formation pathways in China during 2014,
- 1733 Atmospheric Environment, 213, 133-147, 2019.
- 1734 <u>Yumimoto, K., Uno, I., Sugimoto, N., Shimizu, A., Liu, Z., and Winker, D. M.: Adjoint inversion</u>
- modeling of Asian dust emission using lidar observations, Atmospheric Chemistry and Physics,
- 1736 <u>8, 2869-2884, 2008.</u>
- 2737 Zhang, F., Weng, Y., Sippel, J. A., Meng, Z., and Bishop, C. H.: Cloud-Resolving Hurricane
- 1738 Initialization and Prediction through Assimilation of Doppler Radar Observations with an
- 1739 Ensemble Kalman Filter, Monthly Weather Review, 137, 2105-2125, 10.1175/2009mwr2645.1,
- 1740 2009a.
- 1741 Zhang, L., Chen, Y., Zhao, Y., Henze, D. K., Zhu, L., Song, Y., Paulot, F., Liu, X., Pan, Y., Lin, Y.,
- and Huang, B.: Agricultural ammonia emissions in China: reconciling bottom-up and top-down
- estimates, Atmospheric Chemistry and Physics, 18, 339-355, 2018.
- Zhang, Q., Streets, D. G., Carmichael, G. R., He, K. B., Huo, H., Kannari, A., Klimont, Z., Park, I.
- 1745 S., Reddy, S., Fu, J. S., Chen, D., Duan, L., Lei, Y., Wang, L. T., and Yao, Z. L.: Asian emissions
- in 2006 for the NASA INTEX-B mission, Atmospheric Chemistry and Physics, 9, 5131-5153,
- 1747 10.5194/acp-9-5131-2009, 2009b.
- 1748 Zhang, S., Zheng, X., Chen, J. M., Chen, Z., Dan, B., Yi, X., Wang, L., and Wu, G.: A global carbon
- assimilation system using a modified ensemble Kalman filter, Geosci. Model Dev., 8, 805-816,
- 1750 10.5194/gmd-8-805-2015, 2015.
- 1751 Zhang, X., Liu, J., Han, H., Zhang, Y., Jiang, Z., Wang, H., Meng, L., Li, Y. C., and Liu, Y.: Satellite-
- Observed Variations and Trends in Carbon Monoxide over Asia and Their Sensitivities to
- 1753 Biomass Burning, Remote Sensing, 12, 10.3390/rs12050830, 2020.
- 1754 Zheng, B., Tong, D., Li, M., Liu, F., Hong, C., Geng, G., Li, H., Li, X., Peng, L., Qi, J., Yan, L.,
- Zhang, Y., Zhao, H., Zheng, Y., He, K., and Zhang, Q.: Trends in China's anthropogenic
- 1756 emissions since 2010 as the consequence of clean air actions, Atmospheric Chemistry And
- 1/757 Physics, 18, 14095-14111, 10.5194/acp-18-14095-2018, 2018.

1758	Zheng, B., Zhang, Q., Tong, D., Chen, C., Hong, C., Li, M., Geng, G., Lei, Y., Huo, H., and He, K.:
1759	Resolution dependence of uncertainties in gridded emission inventories: a case study in Hebei,
1760	China, Atmospheric Chemistry and Physics, 17, 921-933, 2017.