



1 Simulating  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  in CMAQ version 5.2.1, based on  
2  $^{15}\text{N}$  incorporated SMOKE version 4.6 and WRF version 4.0 for assessing  
3 the role atmospheric processes plays in controlling the isotopic  
4 composition of  $\text{NO}_x$ ,  $\text{NO}_y$ , and atmospheric nitrate

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6 *Huan Fang<sup>†</sup> and Greg Michalski<sup>†‡</sup>*

7 <sup>†</sup>Department of Earth, Atmospheric, and Planetary Sciences Purdue University, 550 Stadium Mall  
8 Drive, West Lafayette, IN 47907, United States

9 <sup>‡</sup>Department of Chemistry, Purdue University, 560 Oval Drive, West Lafayette, IN 47907, United  
10 States

11

12

13 Correspondence: Huan Fang, [fang63@purdue.edu](mailto:fang63@purdue.edu)

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## 1 Abstract

2 Nitrogen oxides ( $\text{NO}_x = \text{nitric oxide (NO)} + \text{nitrogen dioxides (NO}_2)$ ) are important trace gases  
3 that affect atmospheric chemistry, air quality, and climate. Despite the importance of  $\text{NO}_x$   
4 emissions, there are significant uncertainties in  $\text{NO}_x$  emission inventories. After  $\text{NO}_x$  from  
5 different sources being emitted into the atmosphere, its composition will change due to  
6 atmospheric processes. In this study, we used the nitrogen stable isotope composition of  $\text{NO}_x$   
7 ( $\delta^{15}\text{N}(\text{NO}_x)$ ) to trace the changes in  $\delta^{15}\text{N}$  values along the “journey” of atmospheric  $\text{NO}_x$ , by  
8 incorporating  $^{15}\text{N}$  into the emission input dataset prepared from the previous companion research  
9 (Fang & Michalski, 2020) to run CMAQ (the Community Multiscale Air Quality Modeling  
10 System). The simulated spatiotemporal patterns in  $\text{NO}_x$  isotopic composition were compared with  
11 corresponding atmospheric measurements in West Lafayette, Indiana, USA. The results indicate  
12 that estimating of atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  using CMAQ shows better agreement with observation  
13 than using SMOKE (Sparse Matrix Operator Kernel Emissions), due to the consideration of  
14 mixing, disperse, transport, and deposition of  $\text{NO}_x$  emission from different sources.

15

## 16 1. Introduction

17  $\text{NO}_x$  ( $\text{NO}_x = \text{NO} + \text{NO}_2$ ) are important trace gases that affect atmospheric chemistry, air  
18 quality, and climate. The  $\text{NO}_x$  could be converted into  $\text{NO}_y$  ( $\text{NO}_y = \text{NO}_x + \text{HONO} + \text{HNO}_3 +$   
19  $\text{HNO}_4 + \text{N}_2\text{O}_5 + \text{other N oxides}$ ) in the atmospheric  $\text{NO}_x$  cycle. During this process, the ground-  
20 level concentration of  $\text{O}_3$  is elevated and secondary particles are generated. Secondary aerosols in  
21 turn affect cloud physics, enhancing the reflection of solar radiation (Schwartz, S. E., 1996) and  
22 are hazardous to human health (Lighty et al., 2000). Due to its impacts on air quality, climate,  
23 human health, and the environment understanding the spatial and temporal variation in the sources  
24 of  $\text{NO}_x$  is a vital scientific question. However, there are still a number of significant uncertainties  
25 in the  $\text{NO}_x$  budget despite years of research. These include a). soil  $\text{NO}_x$  emissions caused by the  
26 application of N fertilizers (Shepherd, 1991; Ludwig et al., 2001; Galloway et al., 2004; Hudman,  
27 2012; Houlton et al., 2013; Pilegaard, 2013) and the role of vegetation (Johansson, 1987; Jacob &  
28 Wofsy, 1990; Hanson & Lindberg, 1991; Yienger & Levy II, 1995; Thoene, Rennenberg & Weber,  
29 1996; Slovik et al., 1996; Webber & Rennenberg, 1996; Almaraz et al., 2018); b). emissions from  
30 on-road vehicles estimated by different algorithms (Pierson et al., 1996; Singer & Harley, 1996;  
31 Cicero-Fernandez et al., 1997; Dreher & Harley, 1998; Dreher & Harley, 1998; Sawyer et al., 2000;



1 Parrish, 2006); and c). power plant NO<sub>x</sub> emissions due to the implementation of different NO<sub>x</sub>  
2 emission control technologies (Felix et al., 2012; Srivastava et al., 2005; Xing et al., 2013).

3 Previous research has shown that there are distinctive differences in δ<sup>15</sup>N values for NO<sub>x</sub> from  
4 different emission sources (Fig. 1), such as soil (Li & Wang, 2008; Felix & Elliott, 2014; Yu &  
5 Elliott, 2017; Miller et al., 2018), wastes (Felix & Elliott, 2014), vehicles (Moore, 1977; Heaton,  
6 1990; Ammann et al., 1999; Pearson et al., 2000; Savard et al., 2009; Redling et al., 2013; Fibiger,  
7 2014; Felix & Elliott, 2014; Walters et al., 2015a; Walters et al., 2015b), and power plants (Heaton,  
8 1987; Heaton, 1990; Snape, 2003; Felix et al., 2012; Felix et al., 2015; Walters et al., 2015a; Savard  
9 et al., 2017). Thus, the nitrogen stable isotope composition (δ<sup>15</sup>N) of NO<sub>x</sub> could be an effective  
10 tracer of atmospheric NO<sub>x</sub> sources. The δ<sup>15</sup>N(NO<sub>x</sub>) is determined by

$$\delta^{15}\text{N}(\text{NO}_x) (\text{‰}) = \left[ \frac{{}^{15}\text{NO}_x / {}^{14}\text{NO}_x}{{}^{15}\text{N}_2 / {}^{14}\text{N}_2}_{\text{air}} - 1 \right] \times 1000 \quad \text{Eq. (1)}$$

12 where <sup>15</sup>NO<sub>x</sub>/<sup>14</sup>NO<sub>x</sub> is the measurement of relative abundance of <sup>15</sup>N to <sup>14</sup>N in atmospheric NO<sub>x</sub>,  
13 compared with the ratio of nitrogen in the air, of which has a <sup>15</sup>N<sub>2</sub>/<sup>14</sup>N<sub>2</sub> = 0.0036.

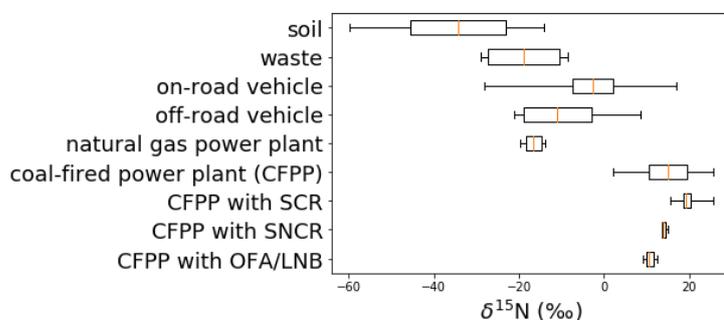


Figure 1: Box (lower quartile, median, upper quartile) and whisker (lower extreme, upper extreme) plot of the distribution of δ<sup>15</sup>N values for NO<sub>x</sub> emission sources

14 Here we have simulated the δ<sup>15</sup>N values of atmospheric NO<sub>x</sub> within the Midwestern United  
15 States, under different scenarios, and compared with the recent measurements. The factors required  
16 to account for the processes that alter δ<sup>15</sup>N of atmospheric NO<sub>x</sub> during the NO<sub>x</sub> chemical lifetime  
17 are: a). The variability of the δ<sup>15</sup>N values of NO<sub>x</sub> emissions in time and space; b). The transport  
18 and mixing of tropospheric NO<sub>x</sub> by meteorology; c.) The wet and dry deposition of NO<sub>x</sub>/NO<sub>y</sub>; and  
19 d). The isotope effects occurring during the tropospheric photochemistry that transforms NO<sub>x</sub> into  
20 NO<sub>y</sub>. In a companion paper (Fang & Michalski, 2020), we discussed the effects due to the variation  
21 of the δ<sup>15</sup>N value of different NO<sub>x</sub> emission sources and their variation in time and space. In this



1 previous study (Fang & Michalski, 2020),  $^{15}\text{N}$  was incorporated into the US EPA trace gas  
2 emission model SMOKE (Sparse Matrix Operator Kernel Emissions), in order to simulate the  
3 spatiotemporal patterns in the isotopic composition  $\text{NO}_x$  and compare them with corresponding  
4 atmospheric measurements. However, the variability in  $\text{NO}_x$  emissions over time and space is not  
5 sufficient to resolve the spatial and temporal changes in the measured  $\delta^{15}\text{N}$  values, due to the bias  
6 of the static SMOKE output files. For example,  $\text{NO}_x$  emitted from a single grid cell dominated by  
7 a coal-fired power plant would result in a  $\text{NO}_x$   $\delta^{15}\text{N}$  around +12‰. If this grid cell were surrounded  
8 by a large array of grid cells dominated by agricultural land-use with a  $\delta^{15}\text{N}$  of -30‰, the  
9 impression is that the region would have a  $\delta^{15}\text{N}$  value close to -30‰. However, since the power  
10 plant emits much more  $\text{NO}_x$  than the surrounding agricultural fields what would be expected for  
11 the actual regional  $\text{NO}_x$   $\delta^{15}\text{N}$  value is biased by the finer emission grid scale. This bias was reduced  
12 by weighting each grid cell's  $\text{NO}_x$  emission relative to the regional total (Fang & Michalski, 2020).  
13 In other words, the fine-scale grids were expanded to larger grids, based on assumptions about the  
14  $\text{NO}_x$  lifetime and transport length scales. This is an unsatisfactory method since the transport of  
15 atmospheric  $\text{NO}_x$  is not controlled by radial diffusion, rather by meteorology/eddy diffusion driven  
16 by pressure gradients.

17 In this work, we explore the effects from the second and third factors, the impacts from  
18 atmospheric transport and deposition processes, by incorporating an input dataset of  $^{15}\text{N}$  emissions  
19 used in simulations by the Chemistry-Transport Model (CTM) used in CMAQ (The Community  
20 Multiscale Air Quality Modeling System). We have previously explored the isotope effects arising  
21 from tropospheric photochemistry using a 0D box model (Michalski et al., 2020). This  $^{15}\text{N}$  isotope  
22 reaction scheme will be incorporated into CMAQ as a new chemical mechanism in order to use  
23 CMAQ to simulate the  $\delta^{15}\text{N}$  of  $\text{NO}_y$  compounds in the subsequent research. The goal of this paper  
24 is to explore how atmospheric processes alter the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  in time and space in  
25 the Midwestern US in the absence of isotope effects occurring during the photochemical  
26 transformation of  $\text{NO}_x$  (source and mixing hypothesis).

27

## 28 2. Methodology

29 In this study, we investigate the role of meteorological transport and removal processes play  
30 in the spatiotemporal distribution of  $\text{NO}_x$   $\delta^{15}\text{N}$  values. The  $^{15}\text{N}$  emission dataset previously  
31 developed (Fang & Michalski, 2020) was used as input for CMAQ to simulate the meteorological  
32 transport effects (advection, eddy diffusion, etc). In addition, CMAQ simulated the effect of  $\text{NO}_x$



1 (nitrate) removal by dry and wet deposition rate is assessed to determine the role of chemistry and  
2 deposition might play in the  $\delta^{15}\text{N}$  of  $\text{NO}_x$  and atmospheric nitrate. The isotope effects associated  
3 with the photochemical transformation of  $\text{NO}_x$  into  $\text{HNO}_3$  and other higher N oxides are ignored,  
4 therefore, this paper only focuses on mixing effects and “lifetime chemistry”, which blur the grid  
5 specific  $\text{NO}_x$   $\delta^{15}\text{N}$  value across the regional scale. The simulations using the same 2002 National  
6 Emission Inventory (NEI) but different meteorology conditions (2002 and 2016) were compared,  
7 in order to explore how meteorology condition impacts the atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$ . Then  
8 simulations using the same meteorology condition (2016) but different emission inventories (2002  
9 NEI and 2016 NEI) were compared, in order to explore how emission inventory impacts the  
10 atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$ . The simulations cover the full domain and nested domain were conducted,  
11 in order to explore and eliminate the bias near the domain boundary.

12

## 13 2.1 The domain of the study

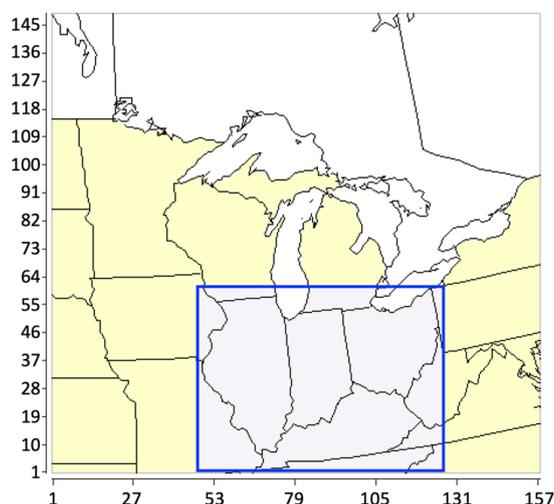


Figure 2: The full geographic domain (yellow) and nested domain (light purple) for the study.

14

15 Two domains were used, a larger domain encompassing the Midwestern region of the United  
16 States and a smaller nested domain of the central portion of the Midwest domain (Fig. 2). The  
17 larger Midwestern domain coordinates ranged from 37 N to 45 N in latitudes, and 98 W to 78 W  
18 in longitude. This fully covers the states of Minnesota, Iowa, Missouri, Wisconsin, Illinois,  
19 Michigan, Indiana, Kentucky, Ohio, and West Virginia, and partially covers North Dakota, South



1 Dakota, Nebraska, Kansas, Tennessee, North Carolina, Virginia, Maryland, Pennsylvania, and  
2 New York (Fig. 2, in yellow). A nested domain, which fully covers the states of Indiana, Illinois,  
3 Ohio, and Kentucky was extracted (Fig. 2, in light purple), in order to reduce the model bias near  
4 the borders that include boundary conditions (details in section 2.6). The horizontal grid resolution  
5 for both domains was 12 km x 12 km. The vertical grid resolution is isobaric dependent, which  
6 increases with height, from 50 m near the surface (bottom layer) to 600 m near the 50 mb pressure  
7 level (top layer).

8

## 9 2.2 $^{15}\text{NO}_x$ and $^{14}\text{NO}_x$ emission input dataset

10 The  $\text{NO}_x$  emission input dataset used by the CTM in CMAQ was prepared, based on the steps  
11 described by Fang & Michalski (2020), and are briefly discussed below. The EPA SMOKE model  
12 was used to simulate  $^{14}\text{NO}_x$  emissions. SMOKE converts the annual  $\text{NO}_x$  emission from county-  
13 level emission data contained in the NEI, into hourly emissions and partitions the emission into  
14 the gridded format. The emission input datasets were prepared using both the 2002 and 2016  
15 versions of the NEI. The main sources of  $\text{NO}_x$  emissions in the NEI's are on-road gasoline, on-  
16 road diesel, off-road gasoline, off-road diesel, coal-fired power plant, natural gas power plant, soil,  
17 and livestock wastes, were categorized into four SMOKE processing categories: Biogenic, Mobile,  
18 Point, and Area (Table 1).

19 The  $^{15}\text{N}$  was incorporated into SMOKE outputs, based on  $\text{NO}_x$  emissions from NEI emission  
20 sectors and the corresponding  $\delta^{15}\text{N}$  values previously discussed (Fang & Michalski, 2020; Table  
21 1). The  $^{15}\text{NO}_x$  emitted by each SMOKE processing category was calculated based on the definition  
22 of  $\delta^{15}\text{N}$  (‰).

$$23 \quad ^{15}\text{NO}_x(i) = ^{14}\text{NO}_x(i) \times ^{15}R_{\text{NO}_x}(i) \quad \text{Eq. (2)}$$

24 where  $^{14}\text{NO}_x(i)$  is the  $\text{NO}_x$  emissions for each category (i) obtained from NEI and  $^{15}R_{\text{NO}_x}$  is a  $^{15}\text{N}$   
25 emission factor ( $^{15}\text{NO}_x/^{14}\text{NO}_x$ ) calculated by:

$$26 \quad ^{15}R_{\text{NO}_x}(i) = \left( \frac{\delta^{15}\text{N}_{\text{NO}_x(i)}}{1000} + 1 \right) \times 0.0036 \quad \text{Eq. (3)}$$

27  $\delta^{15}\text{N}_{\text{NO}_x(i)}$  is the  $\delta^{15}\text{N}$  value of each  $\text{NO}_x$  source category (i = area, biogenic, mobile, and point)  
28 (Table 1) and 0.0036 is the  $^{15}\text{N}/^{14}\text{N}$  of air  $\text{N}_2$ , the reference point for  $\delta^{15}\text{N}$  measurements. The  $\delta^{15}\text{N}$   
29 of total  $\text{NO}_x$  emission was calculated by

$$30 \quad \delta^{15}\text{N}_{\text{NO}_x(\text{total})} = \left( \frac{^{15}\text{NO}_x(\text{area}) + ^{15}\text{NO}_x(\text{biog}) + ^{15}\text{NO}_x(\text{mobile}) + ^{15}\text{NO}_x(\text{point})}{^{14}\text{NO}_x(\text{area}) + ^{14}\text{NO}_x(\text{biog}) + ^{14}\text{NO}_x(\text{mobile}) + ^{14}\text{NO}_x(\text{point})} - 1 \right) \times 1000 \quad \text{Eq. (4)}$$



1 where

$$\begin{aligned}
 2 \quad {}^{15}\text{NO}_x(\text{area}) &= \left( \frac{\delta^{15}\text{N}_{\text{NO}_x(\text{waste})}}{1000} + 1 \right) \times 0.0036 \times {}^{14}\text{NO}_x(\text{waste}) \\
 3 \quad &+ \left( \frac{\delta^{15}\text{N}_{\text{NO}_x(\text{off-road gas})}}{1000} + 1 \right) \times 0.0036 \times {}^{14}\text{NO}_x(\text{off-road gas}) \\
 4 \quad &+ \left( \frac{\delta^{15}\text{N}_{\text{NO}_x(\text{off-road diesel})}}{1000} + 1 \right) \times 0.0036 \times {}^{14}\text{NO}_x(\text{off-road diesel}) \quad \text{Eq. (5)}
 \end{aligned}$$

$$\begin{aligned}
 5 \quad {}^{15}\text{NO}_x(\text{mobile}) &= \left( \frac{\delta^{15}\text{N}_{\text{NO}_x(\text{on-road gas})}}{1000} + 1 \right) \times 0.0036 \times {}^{14}\text{NO}_x(\text{on-road gas}) \\
 6 \quad &+ \left( \frac{\delta^{15}\text{N}_{\text{NO}_x(\text{on-road diesel})}}{1000} + 1 \right) \times 0.0036 \times {}^{14}\text{NO}_x(\text{on-road diesel}) \quad \text{Eq. (6)}
 \end{aligned}$$

7 where  $\delta^{15}\text{N}_{\text{NO}_x(\text{on-road gas})} = -12.35 + 3.02 \times \ln(t + 0.455)$

8 Biogenic is  $\text{NO}_x$  emission from by-products of microbial nitrification and denitrification  
 9 occurring in the soil; mobile is  $\text{NO}_x$  emission from the on-road vehicle; the point is  $\text{NO}_x$  emission  
 10 from power plants or industry; and the area is all other stationary anthropogenic  $\text{NO}_x$  emissions,  
 11 which spread over a spatial extent and individually too small in magnitude to report as point source  
 12 These include off-road vehicles (utility vehicles for agricultural and residential purposes),  
 13 residential combustion, industrial processes, agriculture production (livestock waste, fertilizer, etc),  
 14 etc. Using Eq. (2-6) and  $\delta^{15}\text{N}$  values from previous research (Table 1),  ${}^{15}\text{NO}_x$  emission files were  
 15 generated from the SMOKE  ${}^{14}\text{NO}_x$  output files. The  $\delta^{15}\text{N}$  of on-road gasoline vehicles was based  
 16 on the average vehicle travel time (t) within each region with the same zip code (Walters et al.,  
 17 2015a). The average  $\delta^{15}\text{N}$  of on-road gasoline vehicles within the study area is  $-2.7 \pm 0.8\%$ .

18

SMOKE Category	NEI Sector	$\delta^{15}\text{N}\text{-NO}_x$ (‰) in this study
Biogenic	Soil	-34.3 (Felix & Elliott, 2014)
Area	Livestock Waste	-18.8 (Felix & Elliott, 2014)
	Off-road Gasoline	-11.5 (Walters et al., 2015b)
	Off-road Diesel	-10.5 (Walters et al., 2015b)
Mobile	On-road Gasoline	-2.7 (Walters et al., 2015b)
	On-road Diesel	-2.5 (Walters et al., 2015b)



Point	Coal-fired EGUs	+15 (Felix et al., 2012)
	Natural Gas EGUs	-16.5 (Walters et al., 2015)

1 Table 1:  $\delta^{15}\text{N}$  values for  $\text{NO}_x$  emission sources  
2 by SMOKE processing category and NEI sector

3 Since the isotope effects associated with the photochemical transformation of  $\text{NO}_x$  into  $\text{NO}_y$   
4 are ignored,  $^{15}\text{N}$  was not incorporated into the chemical mechanism of CMAQ for the simulations  
5 of this research. Therefore, the  $^{15}\text{NO}_x$  in the emission input dataset acts as a nonreactive chemical.  
6 Since  $^{14}\text{NO}_x$  will go through and be removed in CMAQ's chemical mechanism, the  $^{14}\text{NO}_x$  in the  
7 emission input dataset was replicated and set as a nonreactive chemical. As a result, the mixing  
8 effects on the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  were explored, through the analysis of the time evolution  
9 of nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  concentrations.

10

### 11 2.3 Meteorology input dataset

12 To explore the impact of atmospheric processes, the meteorology input datasets for the year  
13 2002 and 2016 were prepared and compared. The preparation of the meteorology input datasets  
14 for the simulation using CMAQ CTM (CCTM) requires multiple steps. The first step is to generate  
15 the input for the CTM meteorological model using the NARR (North American Regional  
16 Reanalysis) and NAM (North American Mesoscale Forecast System). Both NARR and NAM  
17 Analyses are regional weather model datasets covering North America and were obtained from the  
18 National Centers for Environmental Information (2019). NARR and NAM were used to convert  
19 the weather observations (every 3 hours for NARR, every 6 hours for NAM Analyses) into gridded  
20 meteorological elements, such as temperature, wind field, and precipitation, with the horizontal  
21 resolution of 12 km, and 34 vertical layers, with the thickness, increases with height, from 50 m  
22 near the surface to 600 m near the 50 mb pressure level. The simulation years were 2002 and 2016  
23 and were selected based on the same timeframe as selected  $\text{NO}_y$   $\delta^{15}\text{N}$  measurements. These include  
24 measurements of  $\delta^{15}\text{N}(\text{NO}_3^-)$  at 8 NADP (National Atmospheric Deposition Program) sites within  
25 Indiana, Illinois, Ohio, and Kentucky in 2001-03, and the direct measurements of  $\delta^{15}\text{N}(\text{NO}_x)$   
26 between July and August 2016 (Mase, 2010; Riha, 2013).

27 The second step was to generate the gridded meteorology files on an hourly basis, using the  
28 Weather Research and Forecasting Model (WRF) using the input files prepared by the NARR and  
29 NAM analyses. To maintain consistency between the  $\text{NO}_x$  emission dataset and the meteorology,  
30 the same coordinate system, spatial domain, and grid size used in the SMOKE model were used



1 in the WRF simulation. The same as the emission dataset, the projection type of WRF output is  
2 Lambert Conformal, with the standard parallel of 33 N and 45 N, the central meridian of 97 W.  
3 The output dataset of WRF has the same spatial domain as the emission dataset with a size of 12  
4 km.

5 The last step is to prepare the CMAQ-ready meteorology input dataset based on WRF outputs,  
6 by running MCIP (the Meteorology-Chemistry Interface Processor), one of the major components  
7 of CMAQ. The MCIP first obtains the necessary parameters (Table S1) from WRF outputs. Then  
8 the MCIP extracts the data of the necessary parameters for the appropriate geographic domain,  
9 which are slightly smaller than the domain of WRF outputs since the cells near the boundary are  
10 inadequate for CMAQ simulation. For example, the geographic domain of WRF outputs for this  
11 research is 159 grids in the east-west direction and 150 grids in the north-south direction. Therefore,  
12 MCIP extracts the WRF outputs into 157 grids in the east-west direction and 148 grids in the north-  
13 south direction, which are exactly the same as the emission input dataset prepared from the  
14 previous companion research (Fang & Michalski, 2020), and are adequate for CMAQ simulation.  
15 After that, MCIP converts the units of the parameters into the units, which are consistent with the  
16 CMAQ simulation. For example, the 10-meter wind is displayed as u (east-west) and v (north-  
17 south) component of wind vector in WRF but is displayed as wind speed and wind direction in  
18 CMAQ. If the parameters, which are necessary for running CMAQ, are not available from the  
19 WRF output, MCIP will diagnose and compute them, such as PBL (planetary boundary layer)  
20 parameters and cloud information (cloud top, cloud base, liquid water content, cloud coverage).  
21 The MCIP also conducts the interpolation and mass-weighted averaging of data, if the grid  
22 resolutions of WRF and CMAQ are different. Finally, MCIP organizes the parameters into seven  
23 netCDF files that embedded with I/O API (input/output applications programming interface): 2-D  
24 time-independent fields at cell centers, 2-D time-independent fields on domain perimeter, 2-D  
25 time-independent fields at cell corners, 2-D time-dependent fields at cell centers, 3-D time-  
26 dependent fields at cell centers, 3-D time-dependent fields on domain perimeter, and 3-D time-  
27 dependent fields at cell corners (Table S2).

28

## 29 2.4 The role of deposition

30 The dry and wet deposition rates of nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  were varied to assess their  
31 role in the spatiotemporal distribution of  $\text{NO}_x$   $\delta^{15}\text{N}$  value. First, the dry and wet deposition rate of  
32  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  was set to zero to test the effect of transport and mixing only. This no-deposition



1 simulation was based on 2002 NEI and 2016 meteorology. Next, the dry and wet deposition rate  
2 of nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  was set equal to the CMAQ default (reactive)  $^{14}\text{NO}_x$  rate in the  
3 simulation under the same scenario as the preliminary simulation. An additional simulation under  
4 the same scenario, with the amplified dry and wet deposition rate, was conducted, to utilize as the  
5 “pseudo tropospheric photochemistry” that removes atmospheric  $\text{NO}_x$ . To determine the  
6 deposition rate of nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$ , the initial concentration of  $\text{NO}_x$  was first  
7 magnified to 20 times of the initial concentration derived from the ASCII vertical profiles to  
8 represent a relatively polluted atmospheric chemical condition. At the same time, the emission  
9 rates of nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  were set to zero, in order to explore the removal of  
10 nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  by deposition. After the multiple tuning trials, the deposition velocity  
11 of nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  was set to 30 times of the deposition rate of reactive  $^{14}\text{NO}_x$ , of  
12 which more than 90% of the nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  were removed in the simulation period  
13 of 2 days. This is in effect the same as simulating the conversion of  $\text{NO}_x$  into  $\text{HNO}_3$ , without any  
14 isotope effect, without having to alter the chemical mechanism to include  $^{15}\text{N}$ . This “pseudo  $\text{HNO}_3$ ”  
15 is then removed by wet/dry deposition and the 2-day criteria is the estimated lifetime of  $\text{NO}_x$  in  
16 the atmosphere. By comparing the CMAQ simulation with different settings of  $\text{NO}_x$  deposition  
17 rate, how the removal of atmospheric  $\text{NO}_x$  by dry and wet deposition impacts the  $\delta^{15}\text{N}$  of  
18 atmospheric  $\text{NO}_x$  was explored.

19

## 20 2.5 Initial condition and boundary condition for the simulation

21 The meteorological fields generated by MCIP were used as the inputs for Initial Conditions  
22 Processor (ICON) and Boundary Conditions Processor (BCON), used for running CCTM of  
23 CMAQ. The ICON program prepares the initial chemical/isotopic concentrations in each of the  
24 3D grid cells for use in the initial time step of the CCTM simulation. For this study, the initial  
25 condition was derived from the ASCII vertical profiles to create a “clean” atmospheric chemical  
26 condition within the domain at the beginning of the simulation, of which the background  
27 concentration of  $\text{NO}_x$  in each grid is lower than 0.25 ppb. The BCON program prepares the  
28 chemical/isotopic boundary condition for throughout the CCTM simulation. Similarly, the  
29 boundary condition was derived from the ASCII vertical profiles for this study, which assume a  
30 “clean” atmospheric chemical condition ( $\text{NO}_x$  concentration lower than 0.25 ppb at surface layer)  
31 outside the domain.



1 The  $^{14}\text{NO}_x$  in the outputs of ICON and BCON were replicated and set as nonreactive chemical.  
2 The same technique was applied to the emission input dataset as well. The nonreactive  $^{15}\text{NO}_x$  were  
3 added to the outputs of ICON and BCON, with the concentration equals to 0.0036 of the  
4 concentrations of reactive  $^{14}\text{NO}_x$ , which assumes  $\delta^{15}\text{N} = 0$  at the initial time step and outside the  
5 domain of the simulation (calculated based on Eq. (1-3)). The nonreactive  $^{14}\text{NO}_x$  and  $^{15}\text{NO}_x$  do not  
6 go through the chemical mechanism within CMAQ so that the effects from tropospheric  
7 photochemistry are excluded, thus only atmospheric processes are explored.

8

## 9 2.6 Different versions of the $\text{NO}_x$ emission inventory

10 The simulated  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  based on different emission inventories varies. In  
11 order to explore how the difference in  $\delta^{15}\text{N}$  of  $\text{NO}_x$  emission impacts the simulated atmospheric  
12  $\text{NO}_x$ , under the same meteorology conditions, and keep the consistency of the simulation at the  
13 same time, two different emission input datasets were prepared. The first dataset was solely based  
14 on NEI-2002. The  $^{15}\text{N}$  was incorporated into the pre-merged SMOKE output, simulated from each  
15 sector of NEI-2002 that contains  $\text{NO}_x$  emission, based on the corresponding  $\delta^{15}\text{N}$  values, before  
16 merging into the emission input dataset for CCTM simulation. The second dataset, as a comparison,  
17 directly obtain the emission rates from the first dataset, except for  $^{15}\text{NO}_x$ . The emission rate of  
18  $^{15}\text{NO}_x$  was determined by the emission rate of  $^{14}\text{NO}_x$ , obtained from the first dataset, and  $\delta^{15}\text{N}$  of  
19 total  $\text{NO}_x$  emission, simulated from  $\text{NO}_x$  emissions from each emission sector based on NEI-2016  
20 and the corresponding  $\delta^{15}\text{N}$  values determined by the previous companion research (Fang &  
21 Michalski, 2020). Thus, the only difference between the two emission input datasets is the  $\delta^{15}\text{N}$  of  
22 the  $\text{NO}_x$  emission over each grid within the domain.

23

## 24 2.7 The simulation over the nested domain

25 As mentioned in section 2.5, atmospheric  $\text{NO}_x$   $\delta^{15}\text{N} = 0\text{‰}$  for initial condition and boundary  
26 condition. As a result, the bias occurs near the border of the research area, mainly under the  
27 following two circumstances: a). When the air mass transports out of the research area (Fig. S1)  
28 since Canada is considered as “emission-free zone”, the atmospheric  $\text{NO}_x$  is diluted, which impacts  
29 its  $\delta^{15}\text{N}$  values, especially for those with extreme  $\delta^{15}\text{N}$  values ( $\delta^{15}\text{N} < -15\text{‰}$  or  $\delta^{15}\text{N} > 5\text{‰}$ ); b).  
30 When the air mass with  $\delta^{15}\text{N}(\text{NO}_x)=0$  transports from the “emission-free zone” to the research area  
31 (Fig. S2), the atmospheric  $\text{NO}_x$  is flattened. Therefore, to avoid the bias near the border, the nested  
32 domain that only covers Indiana, Illinois, Ohio, and Kentucky was determined, where the



1 measurements of  $\delta^{15}\text{N}$  values at NADP sites are available (Mase, 2010; Riha, 2013). The boundary  
2 condition for the simulation over the nested domain is extracted from the CCTM output of the full-  
3 domain simulation (BCON code available on Zenodo.org (10.5281/zenodo.4311986)).  
4  
5



1 3. Results and Discussion

2

3 3.1 Simulated spatial variability in  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$

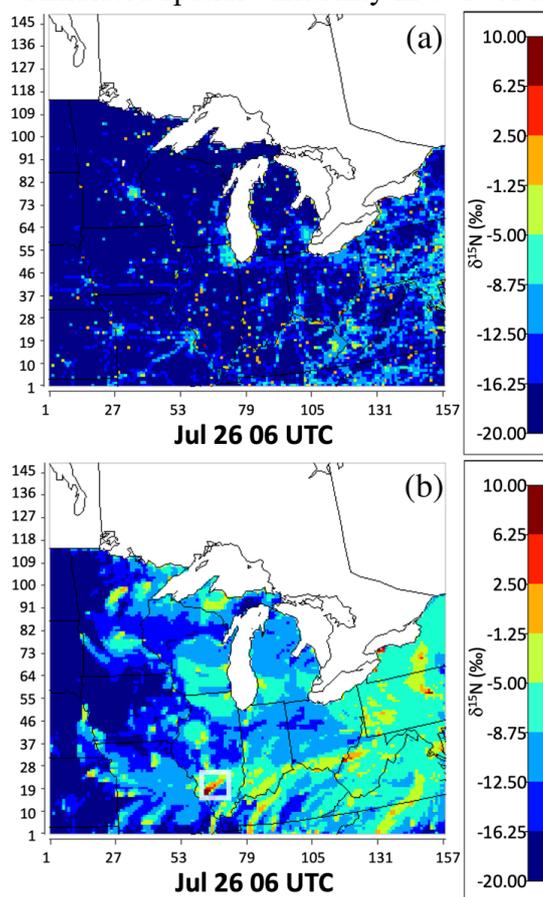


Figure 3: The  $\delta^{15}\text{N}$  values of  $\text{NO}_x$  emission, based on NEI-2002 (a: “no transport” scenario) and the  $\delta^{15}\text{N}$  values of atmospheric  $\text{NO}_x$  based on NEI-2002 and 2016 meteorology (b: “with transport” scenario), at 06 UTC on July 26, are presented by color in each grid. The warmer the color, the higher  $\delta^{15}\text{N}$  values of  $\text{NO}_x$  emission.

4



1 We first examine the spatial heterogeneity of  $\text{NO}_x$   $\delta^{15}\text{N}$  values at a specific time within the  
2 Midwestern domain and explore how atmospheric processes alter the  $\delta^{15}\text{N}$  values relative to the  
3 “no transport” simulation. The “no transport” simulation of  $\text{NO}_x$   $\delta^{15}\text{N}$  values (at 06 UTC on July  
4 26) shows that the domain grids ranged from -34.3‰ to 14.9‰ (Figure 3a). The majority of the  
5 grids within the domain have  $\text{NO}_x$   $\delta^{15}\text{N}$  values lower than -16.3‰. These low  $\delta^{15}\text{N}$  values across  
6 most of the domain are due to the  $\delta^{15}\text{N}$  of -34.3‰ for biogenic  $\text{NO}_x$  emission sources (nitrification  
7 and denitrification) in sparsely populated areas where intensive agriculture dominates the land use  
8 (Fig. 3a). The  $\delta^{15}\text{N}$  values of  $\text{NO}_x$  emitted into grids within big cities mainly ranged between -  
9 8.75‰ and -5‰. This is due to the higher fraction of  $\text{NO}_x$  emission from on-road vehicles having  
10 a  $\delta^{15}\text{N}$  of  $-2.7 \pm 0.8$ ‰. The fraction of  $\text{NO}_x$  emission from on-road vehicles at the grids resolve  
11 major highways is relatively lower, comparing to the grids within big cities, while still higher than  
12 most of the grids within the domain. Thus, the  $\delta^{15}\text{N}$  values along the major highways ranged  
13 between -16.25‰ and -8.75‰. The highest value of  $\delta^{15}\text{N}$  occurs at the grids, where the coal-fired  
14 EGUs (+15‰) and hybrid-fired EGUs (using both coal and natural gas (-16.5‰) for combustion)  
15 are dominant, showing gold (-1.25‰ ~ +2.5‰) and red/dark red (+2.50‰ and above) on the map  
16 (Fig. 3a).

17 The effect of atmospheric mixing and transport on the  $\text{NO}_x$   $\delta^{15}\text{N}$  spatial distribution were then  
18 taken into account by coupling the  $^{15}\text{NO}_x$  emissions (Fang & Michalski, 2020) to the meteorology  
19 simulation. There are significant differences between  $\delta^{15}\text{N}(\text{NO}_x)$  values in the “no transport” (Fig.  
20 3a) and the “with transport” (Fig. 3b) simulations. For example, under the “no transport” scenario  
21 (Fig. 3a) the map of  $\delta^{15}\text{N}(\text{NO}_x)$  values clearly shows the locations of big cities, major highways,  
22 and power plants, but these features are much less obvious in the “with transport” (Fig. 3b)  
23 simulations. The isotopically heavier  $\text{NO}_x$  emission from big cities, such as Chicago, Detroit,  
24 Minneapolis-St Paul, Kansas City, St. Louis, Indianapolis, and Louisville, disperses to the  
25 surrounding rural areas so that the  $\delta^{15}\text{N}(\text{NO}_x)$  values in rural areas are elevated to values similar  
26 to nearby big cities. Similarly, the  $\text{NO}_x$  emitted along major highways is transported to the  
27 surrounding grids, so that the atmospheric  $\text{NO}_x$  at the grids around the major highways become  
28 isotopically heavier relative to the “no transport” scenario. The most obvious and interesting  
29 example is the influence of grids containing coal-fired EGUs on the surrounding region. For  
30 example, the EGU located in the southwestern border of the state of Illinois, Baldwin Energy  
31 Complex (marked with a transparent white box on Fig. 3b), using refined coal, subbituminous coal,  
32 and bituminous coal as its major energy source. The  $\delta^{15}\text{N}(\text{NO}_x)$  in the regions is altered as a



1 function of distance away from the EGU and in this time snapshot, the northeastwards propagating  
2 plume of NO<sub>x</sub> emission from the EGU creates higher  $\delta^{15}\text{N}(\text{NO}_x)$  over 103 km away. Overall, the  
3 “with transport”  $\delta^{15}\text{N}(\text{NO}_x)$  map is indicating the emission source that impacts each grid the most,  
4 after taking atmospheric mixing and transport into account. The domain average  $\delta^{15}\text{N}$  increases  
5 from -20.23‰ under the “no transport” scenario to -11.49‰ under the “with transport” scenario.  
6 The overall emission pattern of the  $\delta^{15}\text{N}$  value shows that the biogenic emission dominates the  
7 spatial domain but after considering the atmospheric processes, anthropogenic emission, mainly  
8 from on-road vehicles, becomes dominant over most of the grids, especially for the grids located  
9 in the suburb of major cities.

10

11 3.2 Seasonal variation in  $\delta^{15}\text{N}$  of NO<sub>x</sub>

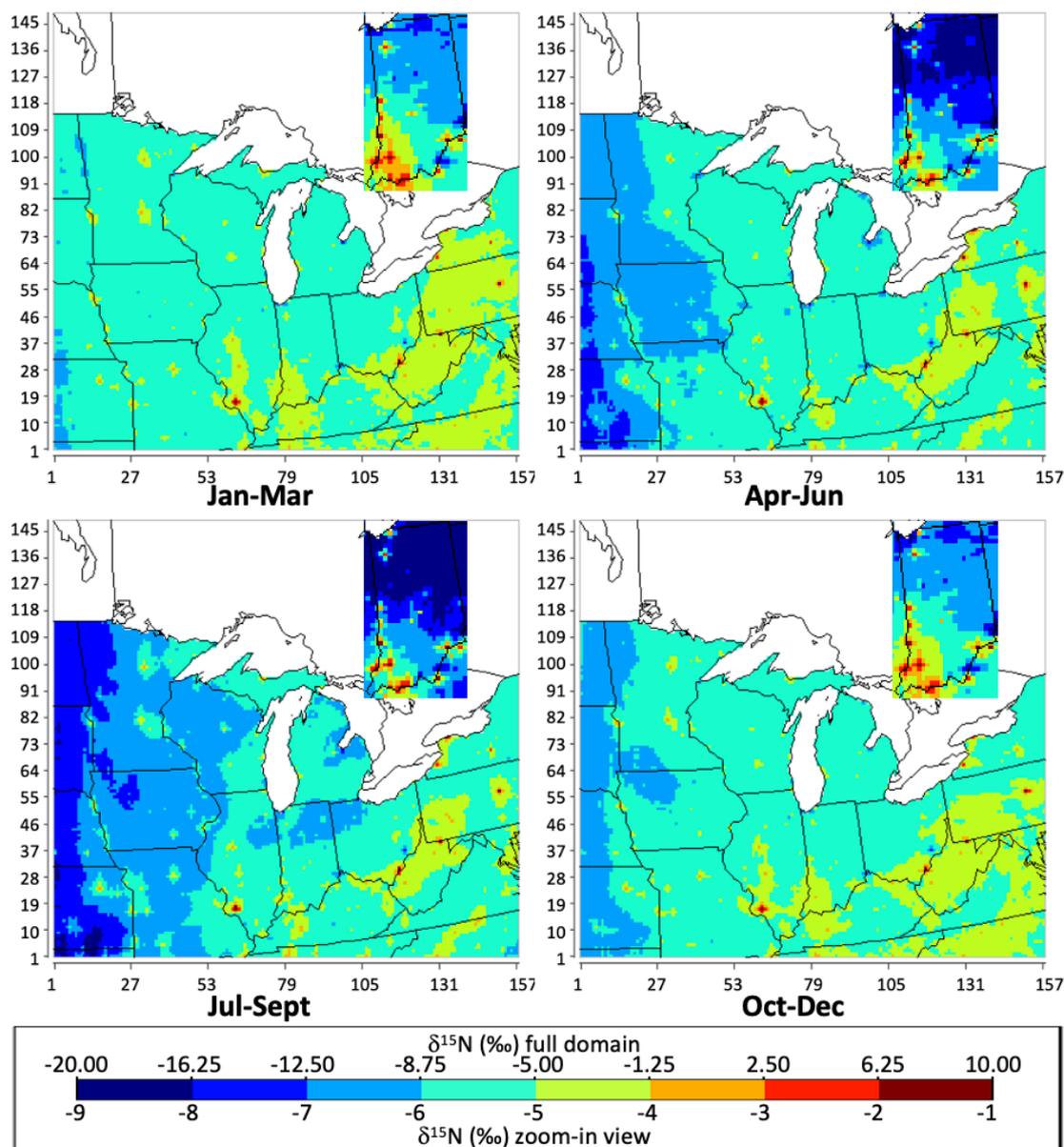


Figure 4: The geographical distribution of the  $\delta^{15}\text{N}$  value of atmospheric  $\text{NO}_x$  in each season (Winter: Jan-Mar; Spring: Apr-Jun; Summer: Jul-Sep; Fall: Oct-Dec) in per mil (‰) throughout the Midwest (with zoom-in view focusing on Indiana) simulated by CMAQ, based on NEI-2002 and 2016 meteorology.

1 We next examine the temporal heterogeneity of atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  under the “with  
2 transport” scenario over the domain and interpret them in terms of changes if the propagation of



1 NO<sub>x</sub> emission as a function of time. The predicted seasonal average  $\delta^{15}\text{N}(\text{NO}_x)$  in the Midwest  
2 shows significant variations (Fig. 4). On an annual basis, the  $\delta^{15}\text{N}$  values of NO<sub>x</sub> range from -19.2‰  
3 to 11.6‰, with the annual average over the Midwest domain of -6.10‰, under the “with transport”  
4 scenario. Compared with the seasonal  $\delta^{15}\text{N}(\text{NO}_x)$  under the “no transport” scenario (Fang &  
5 Michalski, 2020), the  $\delta^{15}\text{N}(\text{NO}_x)$  under the “with transport” scenario has a similar overall average  
6 while narrower range, due to the transport and mixing of the air mass driven by the atmospheric  
7 processes. This could be clearly shown on the map, of which the color scheme is smoother,  
8 comparing with the seasonal  $\delta^{15}\text{N}(\text{NO}_x)$  under the “no transport” scenario (Fig. S3). The maps for  
9 different seasons show the obvious changes in  $\delta^{15}\text{N}$  values over western regions of the Midwest,  
10 from -8.75 ~ -5‰ in Oct-Mar to -16.25 ~ -12.5‰ in Apr-Oct.

11 In addition to the variability of the NO<sub>x</sub> emission source, which has been discussed in depth  
12 in the previous companion paper (Fang & Michalski, 2020), the significant temporal variation in  
13 the  $\delta^{15}\text{N}$  value of atmospheric NO<sub>x</sub> during different seasons is controlled by the transport and  
14 mixing of the air mass, under the different meteorology conditions that vary by season. The PBL  
15 height is an effective indicator showing whether the pollutant is under the synoptic condition,  
16 which is favorable for the disperse, mixing, and transport after being emitted into the atmosphere  
17 (Oke, 2002; Shu et al., 2017; Liao et al., 2018; Miao et al., 2019). In order to qualitatively analyze  
18 the changes in  $\delta^{15}\text{N}$  values driven by atmospheric processes, the difference between the  $\delta^{15}\text{N}$  value  
19 of atmospheric NO<sub>x</sub> under the “with transport” scenario and “no transport” scenario ( $\Delta\delta^{15}\text{N}_{\text{transport}}$ )  
20 on the seasonal basis were shown (Fig. 5). The seasonal  $\Delta\delta^{15}\text{N}_{\text{transport}}$  values range from -21.95‰  
21 to 31.22‰, with an average of 4.93‰. The overall pattern of the  $\Delta\delta^{15}\text{N}_{\text{transport}}$  values shows that  
22 after the NO<sub>x</sub> being emitted into the atmosphere, it became isotopically heavier over the majority  
23 of the grids within the domain, and isotopically lighter over the grids that contain big cities, major  
24 highways, and power plants. This could be explained by the transport and disperse of biogenic  
25 emission and anthropogenic emission to the surrounding areas. Among the grids located in rural  
26 areas, where the biogenic emission dominates the NO<sub>x</sub> budget, the  $\delta^{15}\text{N}$  values increases from  
27 around -30‰ to around -10‰, due to transport and disperse of anthropogenic emission with  
28 relatively high emission rates from surrounding cities, highways, or power plants, which brings  
29 the isotopically heavier NO<sub>x</sub> into the grids. On the other hand, among the grids located in the urban  
30 area, highways, or power plants, where anthropogenic emission dominates the NO<sub>x</sub> budget, the  
31 changes in  $\delta^{15}\text{N}$  values decrease is much less obvious, showing the  $\Delta\delta^{15}\text{N}_{\text{transport}}$  values ranges  
32 between -5‰ and +5‰. This could be explained by the relatively high rates of anthropogenic



1 emission. Thus, the effects of the transport and disperse of biogenic emissions from the  
2 surrounding rural area are minimal.

3 Comparing the distributions of the difference in  $\delta^{15}\text{N}$  values (Fig. 5) with the corresponding  
4 PBL height (Fig. S4) among the maps of each season, the effects of PBL height on the propagation  
5 of the air mass are clearly shown. The PBL height changes significantly among each season within  
6 the geographic domain, especially over Minnesota, Wisconsin, and Iowa (Fig. S4). The PBL height  
7 over these area increases from less than 250 meters above the ground level to more than 625 meters  
8 above the ground level, during spring (Apr-Jun) and summer (Jul-Sep), which creates a more  
9 favorable synoptic condition for the disperse, mixing, and transport of the pollutant after being  
10 emitted into the atmosphere. As a result, the difference in  $\delta^{15}\text{N}$  values shifts to higher values,  
11 showing the stronger effect of atmospheric processes during spring and summer. The positive  
12 correlation between PBL height and propagation of air mass, indicated by the evolution of  
13 atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  in this study, agrees well with the corresponding measurement in  
14 megacities in China from the previous studies (Shu et al., 2017; Liu et al., 2018; Liao et al., 2018).

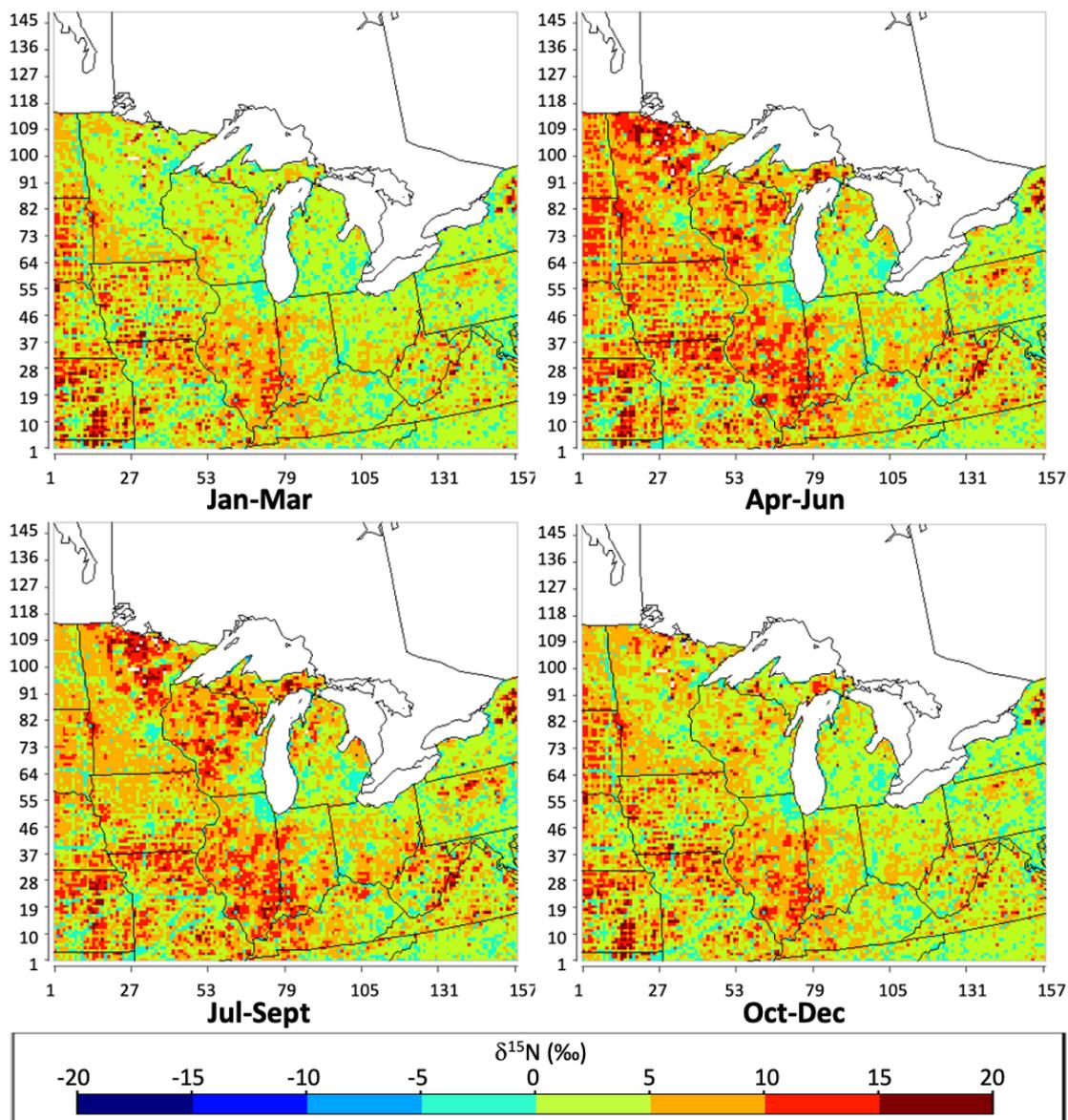


Figure 5: The difference between the  $\delta^{15}\text{N}$  (‰) value of atmospheric  $\text{NO}_x$  under the “with transport” scenario and “no transport” scenario ( $\Delta\delta^{15}\text{N}_{\text{transport}}$ ) during each season (Winter: Jan-Mar; Spring: Apr-Jun; Summer: Jul-Sep; Fall: Oct-Dec), throughout the Midwest simulated by CMAQ, based on NEI-2002 and 2016 meteorology.

1

2



### 1 3.3 Different meteorology conditions

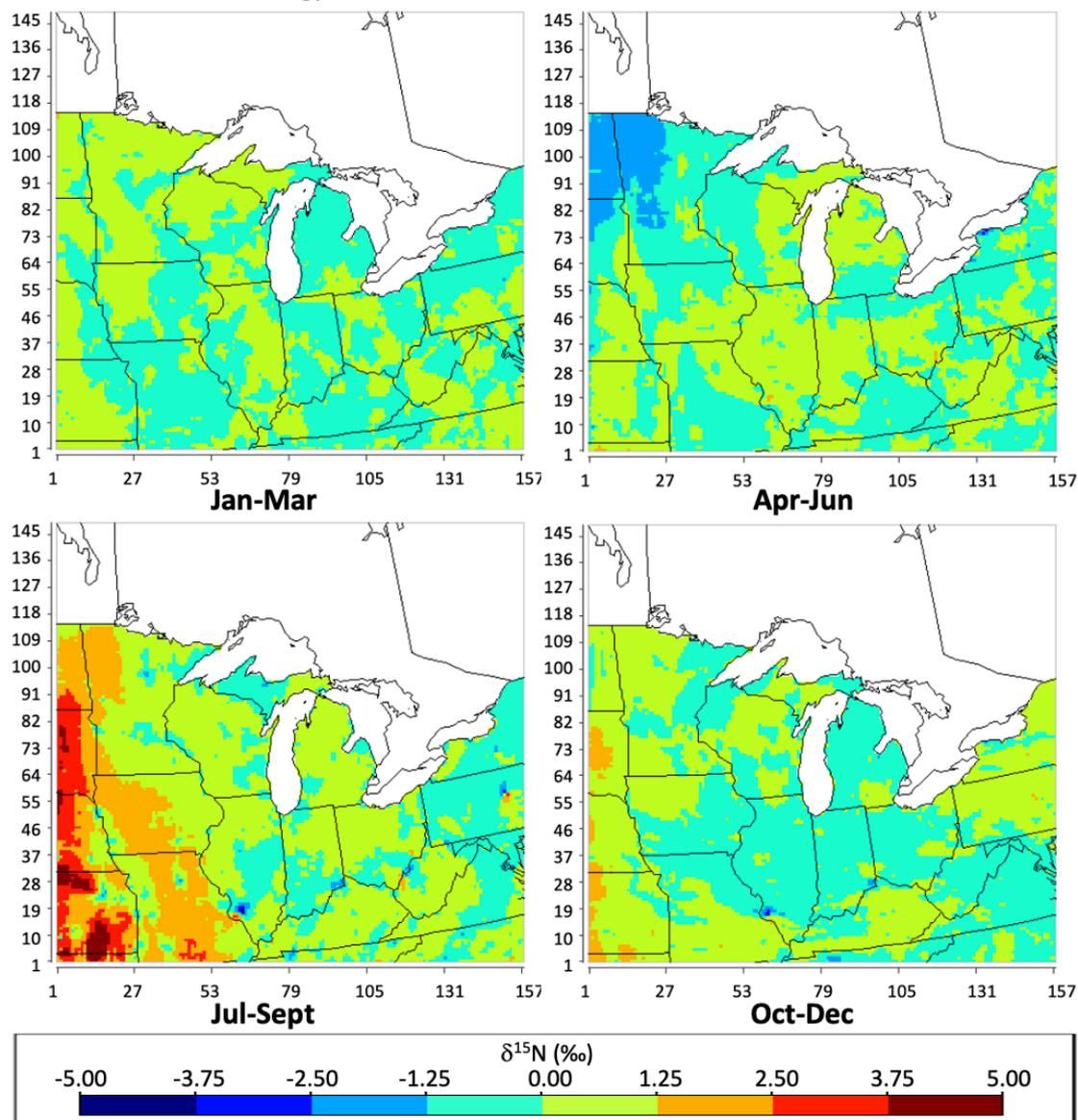


Figure 6: The geographical distribution of the difference between CMAQ simulated  $\delta^{15}\text{N}$  value of atmospheric  $\text{NO}_x$  based on 2002 meteorology and 2016 meteorology in each season (Winter: Jan-Mar; Spring: Apr-Jun; Summer: Jul-Sep; Fall: Oct-Dec) in per mil (‰) throughout the Midwest.

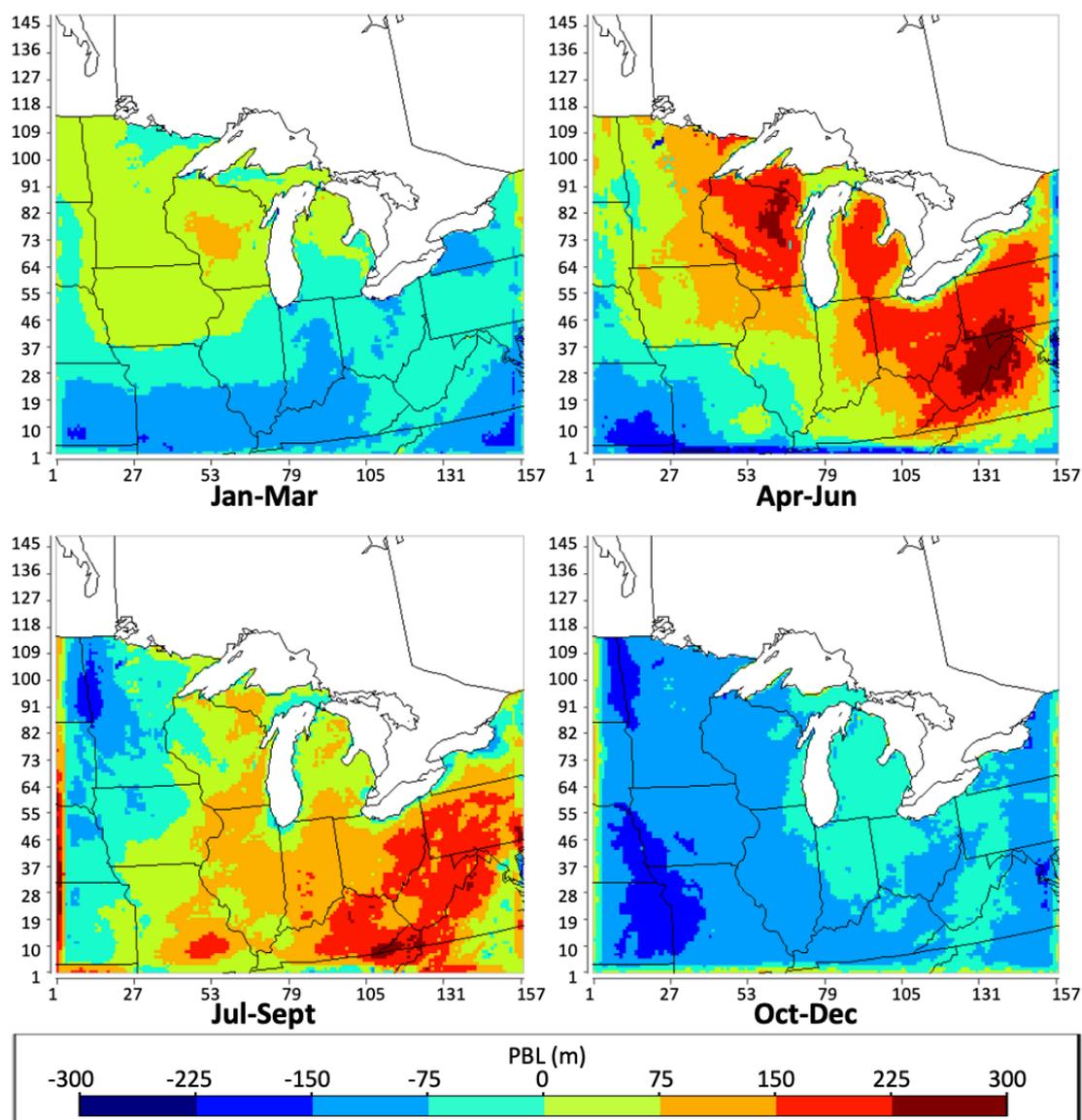


Figure 7: The geographical distribution of the difference between planetary boundary layer (PBL) height in meters based on 2002 meteorology and 2016 meteorology during each season (Winter: Jan-Mar; Spring: Apr-Jun; Summer: Jul-Sep; Fall: Oct-Dec) of 2016 throughout the Midwest.

1 The atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  simulated based on different meteorology input dataset varies. In  
2 order to compare the spatial heterogeneity of the atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  under different  
3 meteorology conditions, the same analysis was done on the simulation using 2002 meteorology



1 (Fig. S6). Overall, the simulated atmospheric  $\text{NO}_x$  under 2002 meteorology is isotopically heavier  
2 than under 2016 meteorology, especially in the western part of the domain during summer (Fig.  
3 6). The dynamics of PBL height potentially cause the variation in the level of disperse, mixing,  
4 and transport of  $\text{NO}_x$  emission. Due to the significantly higher level of PBL during spring and  
5 summer (Fig. 7) comparing to the seasonal PBL height during 2016, the disperse, mixing, and  
6 transport of anthropogenic  $\text{NO}_x$  emission with higher  $\delta^{15}\text{N}$  values alters the atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$   
7 over the rural area further, under 2002 meteorology. The videos of atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  on an  
8 hourly basis throughout the year 2002 and 2016 are available on Zenodo.org  
9 (10.5281/zenodo.4311986).

10

### 11 3.4 Different versions of emission inventories

12 There was a dramatic difference in the atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  simulated depending on which  
13 NEI emission dataset was used. In order to compare the spatial heterogeneity of the atmospheric  
14  $\delta^{15}\text{N}(\text{NO}_x)$  generating from different  $\text{NO}_x$  emission budgets, the same analysis was done on the  
15 simulation using the emission input dataset prepared from the 2016 version of NEI (Fig. S7). In  
16 general, the simulated atmospheric  $\text{NO}_x$  based on 2016 NEI is significantly isotopically lighter  
17 than based on 2002 NEI, especially in the central and eastern parts of the domain (Fig. 8).  
18 According to Fang & Michalski (2020), the fraction of  $\text{NO}_x$  emission from the anthropogenic  
19 source in NEI-2016 was lower than in NEI-2002 for most of the grids within the domain. Therefore,  
20 the atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  based on 2016 NEI was lower. According to US Energy Information  
21 Administration (2017b), from 2002 to 2016, 53 Giga Watts coal-fired and 54 Giga Watts natural  
22 gas EGU retired in the US. The EGU dominates the  $\text{NO}_x$  emission at the grids where it is located  
23 in, account for up to 90% of the total  $\text{NO}_x$  emission (Fang & Michalski, 2020). Given the  $\delta^{15}\text{N}$   
24 value of the  $\text{NO}_x$  emitted from coal-fired EGU is +15‰ (Table 1), the  $\delta^{15}\text{N}$  values of the  
25 atmospheric  $\text{NO}_x$  over the grids that contain the abandoned coal-fired EGU change dramatically  
26 during the period between 2002 and 2016. A similar pattern occurs at the grids that contain the  
27 EGU, which uses both coal and natural gas as its energy source ( $\delta^{15}\text{N} = -0.75\text{‰}$ ). As a result, the  
28 number of “hotspots” with high  $\delta^{15}\text{N}$  values in 2016 is much less than in 2002. While, the change  
29 in  $\delta^{15}\text{N}$  values of the atmospheric  $\text{NO}_x$  over the grids that contain the abandoned natural gas EGU  
30 is not that obvious under the scenario of “with transport” from 2002 to 2016, since the  $\delta^{15}\text{N}$  value  
31 of the  $\text{NO}_x$  emitted from natural gas EGU is -16.5‰ (Table 1), which is similar to the  $\delta^{15}\text{N}$  values  
32 at the surrounding grids after the mixture of  $\text{NO}_x$  emission from biogenic source and on-road



1 vehicles. Besides this, the implementation of NO<sub>x</sub> emission control technologies (SCR, SCNR,  
2 LNB, OFA) decreases the  $\delta^{15}\text{N}$  of the NO<sub>x</sub> emission from power plants and vehicles, thus decrease  
3 the atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$ .

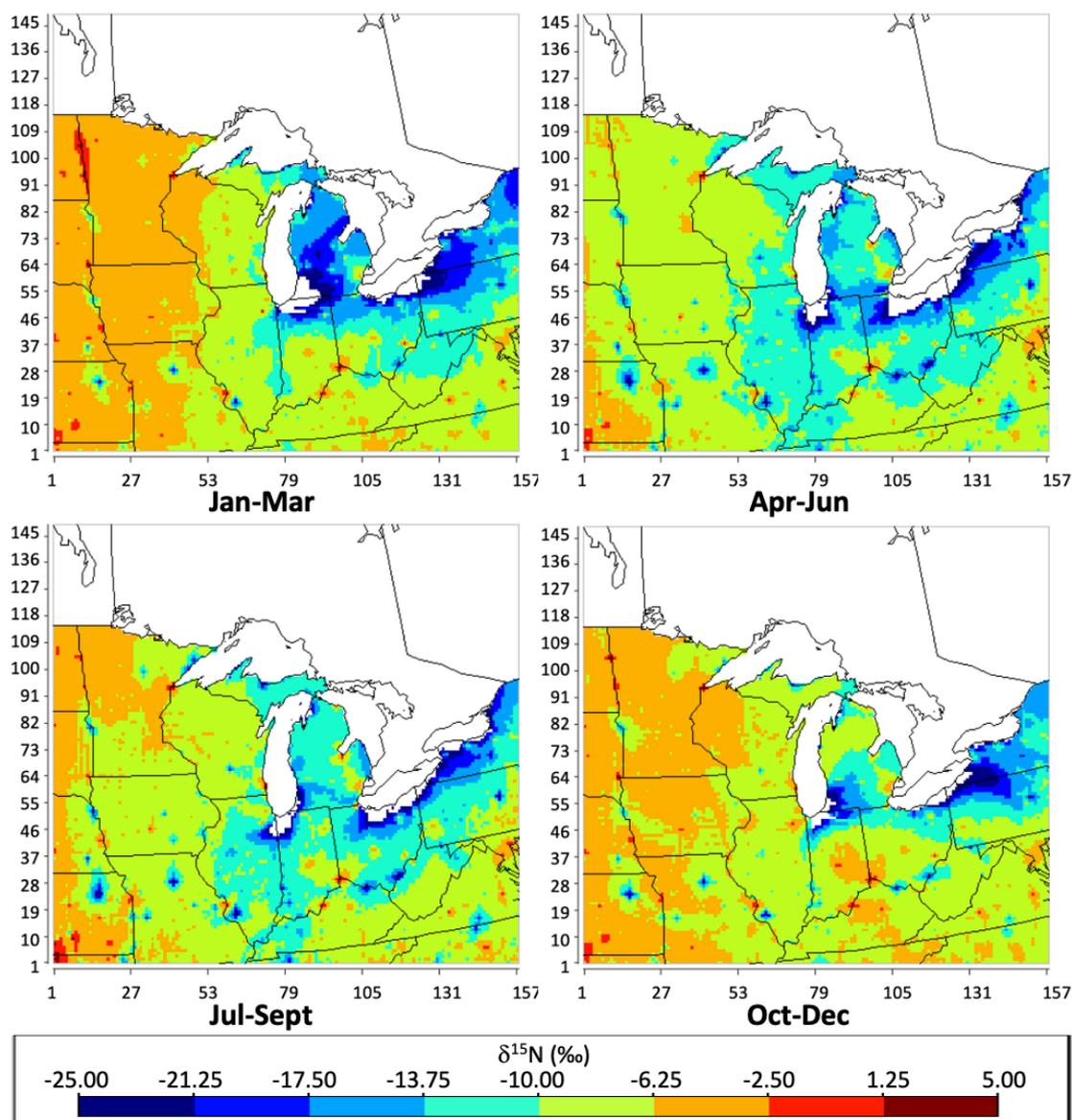


Figure 8: The geographical distribution of the difference between CMAQ simulated  $\delta^{15}\text{N}$  value of atmospheric NO<sub>x</sub> based on NEI-2016 and NEI-2002 in each season (Winter: Jan-Mar; Spring: Apr-Jun; Summer: Jul-Sep; Fall: Oct-Dec) in per mil (‰) throughout the Midwest.



1

## 2 3.5 The role of deposition

3 The deposition alters the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$ . In order to compare the spatial  
4 heterogeneity of the atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  with different settings of  $\text{NO}_x$  deposition rate, the  
5 same analysis was done on the simulation using the amplified dry and wet deposition rates (Fig.  
6 S8). In order to explore the impact of dry and wet deposition on the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$ , the  
7 difference between the  $\delta^{15}\text{N}$  values of atmospheric  $\text{NO}_x$  under the “amplified deposition” scenario  
8 and “default deposition” scenario ( $\Delta\delta^{15}\text{N}_{\text{deposition}}$ ) on the seasonal basis were shown (Fig. 9). The  
9 seasonal  $\Delta\delta^{15}\text{N}_{\text{deposition}}$  values range from  $-3.67\text{‰}$  to  $5.34\text{‰}$ , with an average of  $0.51\text{‰}$ . The overall  
10 pattern of the  $\Delta\delta^{15}\text{N}_{\text{deposition}}$  values shows that due to the impact of deposition, the atmospheric  
11  $\text{NO}_x$  became isotopically lighter over the majority of the grids within the domain, and isotopically  
12 heavier over the grids, which contain or surround power plants and big cities. The amplified  
13 deposition simulation somehow presents the isotope effects associated with the “pseudo  
14 photochemical transformation” of  $\text{NO}_x$  into  $\text{NO}_y$ . The complete isotope effect of tropospheric  
15 photochemistry will be addressed in future work, which incorporates  $^{15}\text{N}$  into the chemical  
16 mechanism of CMAQ for the simulation.

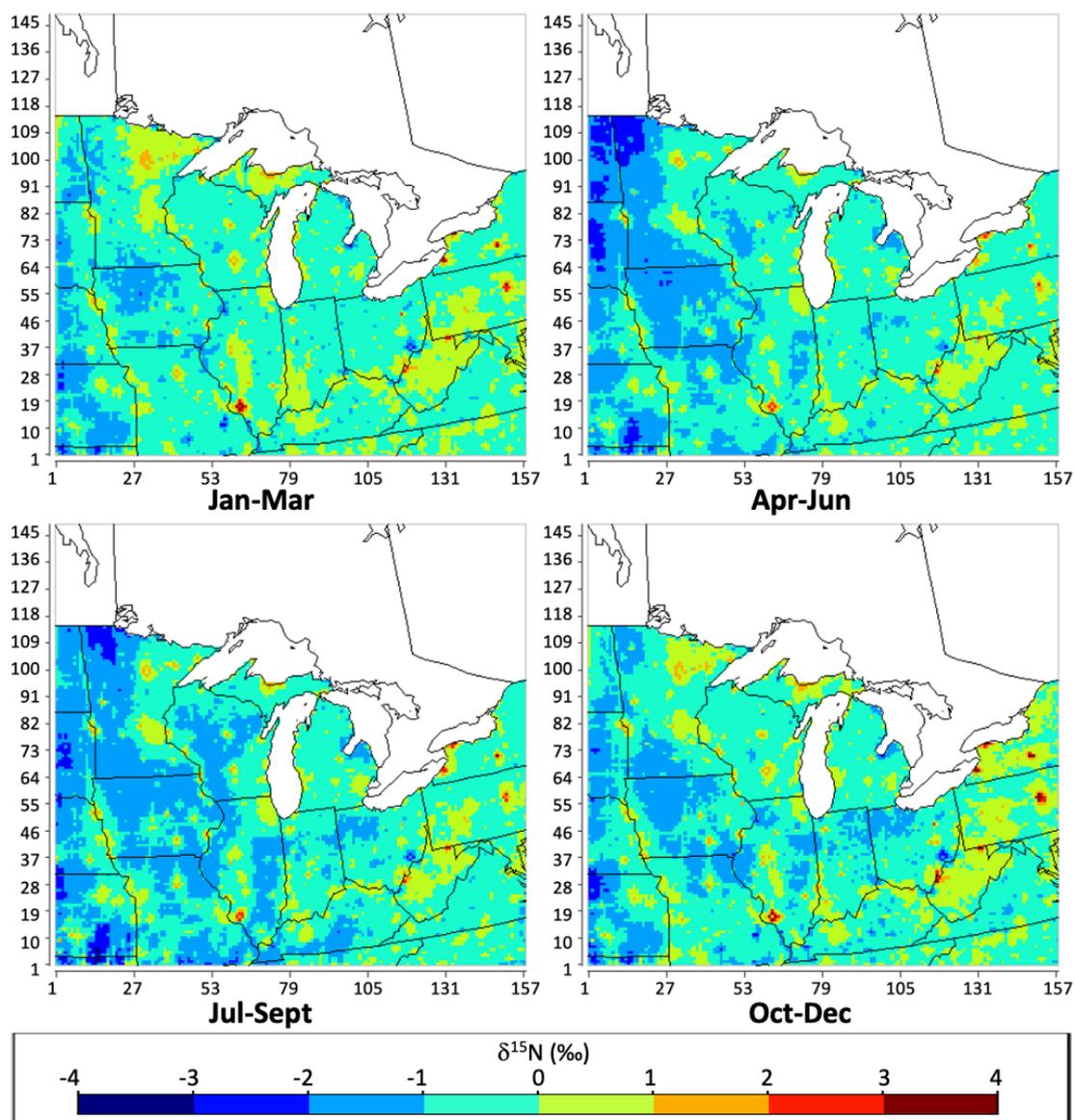


Figure 9: The difference between the  $\delta^{15}\text{N}$  (‰) value of atmospheric  $\text{NO}_x$  under the “amplified deposition” scenario and “default deposition” scenario ( $\Delta\delta^{15}\text{N}_{\text{deposition}}$ ) during each season (Winter: Jan-Mar; Spring: Apr-Jun; Summer: Jul-Sept; Fall: Oct-Dec), throughout the Midwest simulated by CMAQ, based on NEI-2002 and 2016 meteorology.

- 1
- 2 3.6 The simulation over the nested domain



1 We next examine the temporal heterogeneity of difference in atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  between  
2 nested-domain simulation and full-domain simulation ( $\Delta\delta^{15}\text{N}_{\text{nested-full}}$ ), to explore the potential bias  
3 due to the motion of the air mass across the boundary of the geographic domain of the study (Fig.  
4 10). The nested domain covers the states of Indiana, Illinois, Ohio, and Kentucky, where the  
5 measurements of  $\delta^{15}\text{N}$  values at NADP sites are available. The predicted  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$   
6 over the nested domain shows a similar overall pattern as the  $\delta^{15}\text{N}$  within the same domain from  
7 the full-domain simulation, except over the southern border of the domain (Fig. S9). In order to  
8 qualitatively analyze the effects from the initial boundary condition, the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$   
9 within IN, IL, OH, and KY were extracted from the full-domain simulation (Fig. 4) and compare  
10 with the nested-domain simulation within the same region (Fig. 10). The  $\Delta\delta^{15}\text{N}_{\text{nested-full}}$  values  
11 ranged between  $-0.25\%$  and  $+0.25\%$  over most of the grids within the nested domain, showing  
12 the difference between nested-domain simulation and full-domain simulation of  $\delta^{15}\text{N}$  values are  
13 trivial. However, near the southern border of the nested domain, the obvious  $\Delta\delta^{15}\text{N}_{\text{nested-full}}$  values  
14 closed to  $+0.75\%$  during fall and winter, closed to  $+1.00\%$  during spring and summer occur, which  
15 indicate the atmospheric  $\text{NO}_x$  from the nested-domain simulation is isotopically heavier. The  
16 values of  $\Delta\delta^{15}\text{N}_{\text{nested-full}}$  become obvious near the southern border, which indicates the dilution of  
17  $\text{NO}_x$ , after it transports out of the domain since the  $\delta^{15}\text{N}$  on the boundary was set to zero. Unlike  
18 the southern border, the northern, western, and eastern border of the nested domain is located  
19 sufficient distance apart from the boundary of the full domain. As a result, the  $\Delta\delta^{15}\text{N}_{\text{nested-full}}$  values  
20 are similar over the majority grids within the domain.

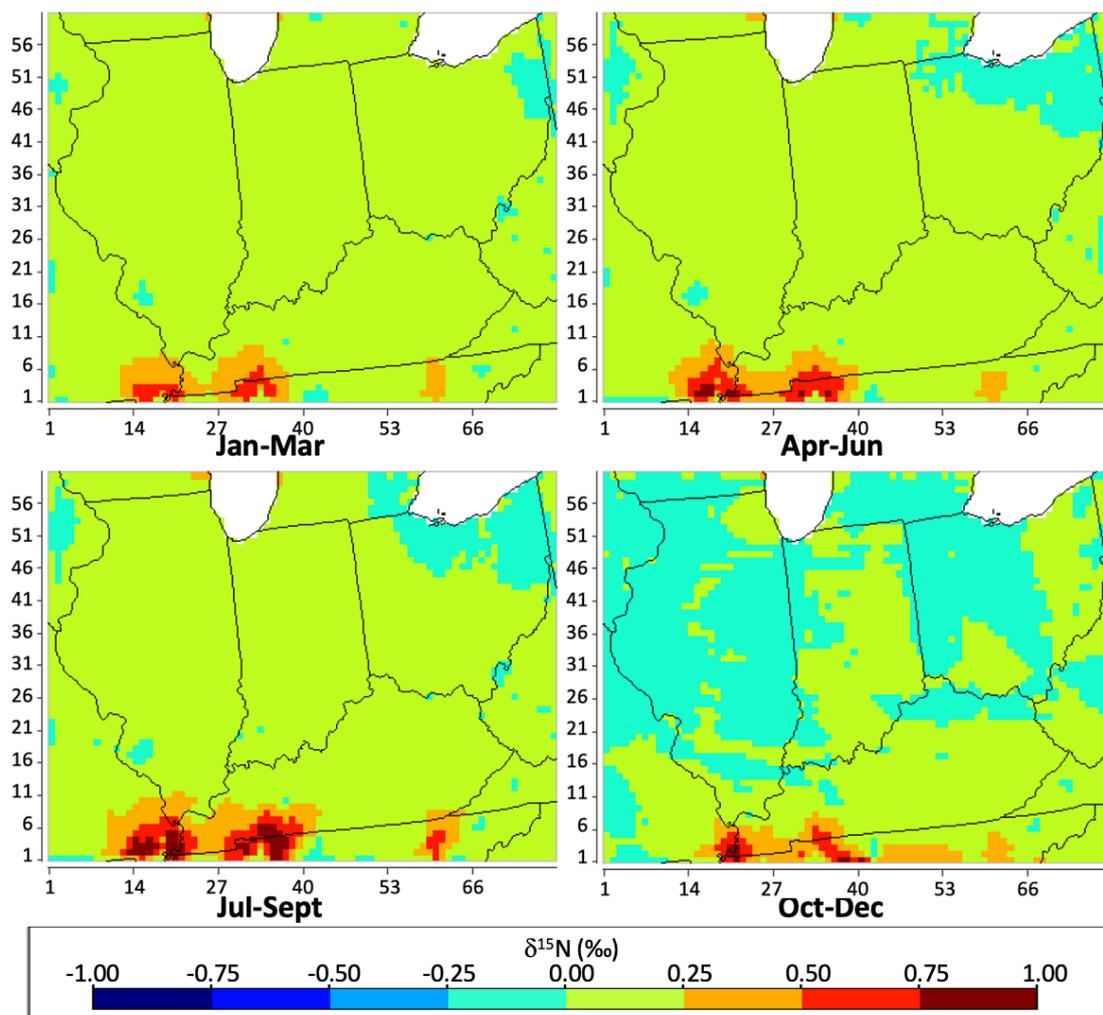


Figure 10: The geographical distribution of the difference between nested-domain simulation and full-domain simulation of  $\delta^{15}\text{N}$  value of atmospheric  $\text{NO}_x$  ( $\Delta\delta^{15}\text{N}_{\text{nested-full}}$ ) in each season (Winter: Jan-Mar; Spring: Apr-Jun; Summer: Jul-Sep; Fall: Oct-Dec) in per mil (‰) within IN, IL, OH, and KY, based on NEI-2002 and 2016 meteorology.

- 1
- 2 3.7 Model-observation comparison

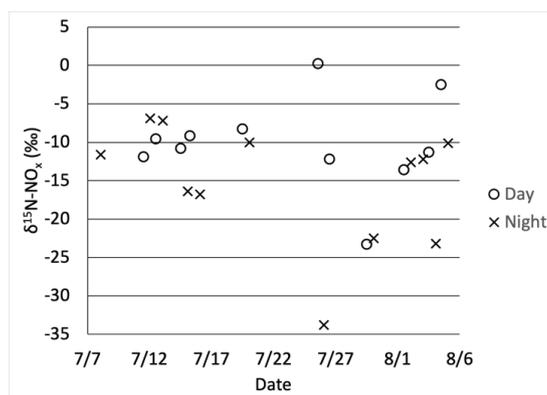


Figure 11: The  $\delta^{15}\text{N}(\text{NO}_x)$  values measured at West Lafayette, IN between July 9 and August 5, 2016, from 8 am to 4 pm during the daytime ( $\circ$ ), and from 9:30 pm to 5:30 am during the nighttime ( $\times$ )

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In order to evaluate the CMAQ simulation of atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$ , several existing datasets of measurements were utilized to compare with the simulations. As the only direct measurements of  $\delta^{15}\text{N}(\text{NO}_x)$  within the domain,  $\text{NO}_x$  samples collected between July 8 and August 5, 2016 (Fig. 11; Walters, Fang, & Michalski, 2018) was first used for the validation of the CMAQ simulation. 30  $\text{NO}_x$  samples were collected from 8 am to 4 pm during the daytime, and from 9:30 pm to 5:30 am during the nighttime in West Lafayette, IN, an NADP (National Atmospheric Deposition Program) site in the northwest part of Indiana and home to Purdue University. The measured  $\delta^{15}\text{N}(\text{NO}_x)$  ranged from -23.3 to 0.2‰ for the daytime samples and -33.8 to -6.9‰ for the nighttime samples.

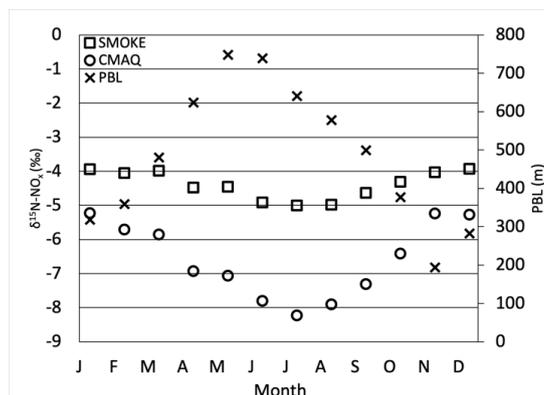


Figure 12: The monthly  $\delta^{15}\text{N}$  values of total  $\text{NO}_x$  emission simulated by SMOKE ( $\square$ ) based on NEI-2002, the monthly  $\delta^{15}\text{N}$  values of atmospheric  $\text{NO}_x$  simulated by CMAQ ( $\circ$ ) based on NEI-2002 and 2016 meteorology, the monthly average of PBL height ( $\times$ , right axis) over the 12-km grid that covers West Lafayette, IN.

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The CMAQ simulated  $\delta^{15}\text{N}$  values of atmospheric  $\text{NO}_x$  in West Lafayette show more obvious monthly variations and seasonal trends comparing to the  $\delta^{15}\text{N}$  values of  $\text{NO}_x$  emission (Fig. 12, in circle ( $\circ$ )). The simulation shows that the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  starts around -5‰ in January, which is about 1‰ lower than  $\delta^{15}\text{N}$  of  $\text{NO}_x$  emission (Fig. 12, in square ( $\square$ )). During winter (Jan-Mar), the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  decrease slightly, and the difference between the  $\delta^{15}\text{N}$  of  $\text{NO}_x$  emission gradually increases. During spring (Apr-June), the more obvious decreasing trend of the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  occurs, and the difference between the  $\delta^{15}\text{N}$  of  $\text{NO}_x$  emission is larger than during winter. The  $\delta^{15}\text{N}$  value reaches the minimum around -8‰ in July. During summer (Jul-Sept), the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  starts to increase, and the difference between the  $\delta^{15}\text{N}$  of  $\text{NO}_x$  emission decreases. During fall (Oct-Dec), the  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  increases, and the difference between the  $\delta^{15}\text{N}$  of  $\text{NO}_x$  emission decreases, but with a slighter trend than during summer. The  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  ends at -5‰, 1‰ lower than  $\delta^{15}\text{N}$  of  $\text{NO}_x$  emission. In addition to the change in the fractions of  $\text{NO}_x$  emission sources from April to September, which has been discussed in the previous companion paper (Fang & Michalski, 2020), the monthly variations and seasonal trend of the simulated atmospheric  $\delta^{15}\text{N}(\text{NO}_x)$  are mainly driven by the



1 strength of disperse, mixing, and transport of the atmospheric NO<sub>x</sub> emitted from different sources,  
 2 indicated by the PBL height. The PBL height during the period from April to September is 90%  
 3 higher than during the period from October to March, which is favorable for the mixture of  
 4 isotopically lighter NO<sub>x</sub> from the surrounding area (Fig. 12, in cross (×)). Thus, the δ<sup>15</sup>N of  
 5 atmospheric NO<sub>x</sub> diverges further from the δ<sup>15</sup>N of NO<sub>x</sub> emission.

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	measurement	NEI 2002	NEI 2016	NEI-2002 +WRF2016	NEI-2002 +WRF2002	NEI-2016 +WRF2016	NEI-2002 +WRF2016 nested
min	-33.800	-12.180	-18.439	-15.824	-14.779	-31.149	-15.858
max	0.200	-3.753	-4.410	-3.360	-3.726	5.458	-3.187
median	-11.250	-4.993	-7.049	-8.094	-8.355	-13.975	-8.108
stdev	8.023	2.168	2.388	2.081	1.881	4.122	2.099

7 Table 2: Performance of δ<sup>15</sup>N(NO<sub>x</sub>) simulation for West Lafayette, IN

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9 The CMAQ simulation of the δ<sup>15</sup>N of atmospheric NO<sub>x</sub> under different scenarios of NEI and  
 10 WRF was compared with the measurement (Walters, Fang, & Michalski, 2018) from July 8 to  
 11 August 5, 2016 (Fig. 13). The δ<sup>15</sup>N of atmospheric NO<sub>x</sub> simulated based on NEI-2002 and 2016  
 12 meteorology ranges from -15.8‰ to -3.4‰, with the medium of -8.1 ± 2.1‰; the δ<sup>15</sup>N of  
 13 atmospheric NO<sub>x</sub> simulated based on NEI-2002 and 2002 meteorology ranges from -14.8‰ to -  
 14 3.7‰, with the medium of -8.4 ± 1.9‰; the δ<sup>15</sup>N of atmospheric NO<sub>x</sub> simulated based on NEI-  
 15 2016 and 2016 meteorology ranges from -31.1‰ to -5.5‰, with the medium of -14.0 ± 4.1‰. The  
 16 δ<sup>15</sup>N of the corresponding measurement ranges from -33.8‰ to 0.2‰, with the medium of -11.2  
 17 ± 8.0‰. In general, the CMAQ simulations of δ<sup>15</sup>N(NO<sub>x</sub>) under most of the scenarios conducted  
 18 in this study, except the simulation based on NEI-2016 and 2016 meteorology, perform better than  
 19 the SMOKE simulation of δ<sup>15</sup>N(NO<sub>x</sub>), which only take the variability of NO<sub>x</sub> emission source into  
 20 account (Table 2). On the other hand, the simulation based on NEI-2016 and 2016 meteorology  
 21 capture the isotopically light NO<sub>x</sub> better than the simulations under the other scenarios of emission  
 22 and meteorology input datasets.

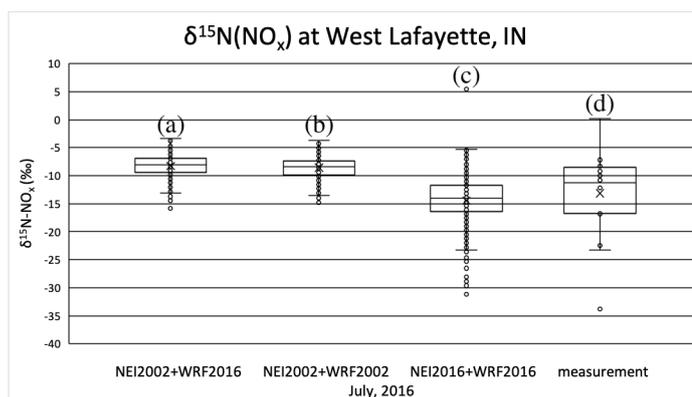


Figure 13: The distributions of  $\delta^{15}\text{N}(\text{NO}_x)$  values over the 12-km grid that covers West Lafayette, IN from July 8 to August 5, simulated by CMAQ, based on NEI-2002 and 2016 meteorology (a), NEI-2002 and 2002 meteorology (b), NEI-2016 and 2016 meteorology (c), compare with the corresponding measurement (d) taken on July to August in 2016 (box: lower quartile, median, upper quartile; whisker: lower extreme, upper extreme; dots outside the whisker: outliers)

1  
2 Finally, we compared the CMAQ predicted  $\delta^{15}\text{N}(\text{NO}_x)$ , under the scenario of NEI-  
3 2002+WRF2002 at NADP sites within Indiana, Illinois, Ohio, and Kentucky (Table S3) with the  
4 measurements of  $\delta^{15}\text{N}(\text{NO}_3^-)$  from 2001 to 2003. The  $\delta^{15}\text{N}$  values of atmospheric  $\text{NO}_x$  simulated  
5 by CMAQ at these sites show obvious monthly variations and seasonal trends (Fig. 14, top). The  
6 monthly boxes are the 1<sup>st</sup> and 3<sup>rd</sup> quantiles of the simulated monthly  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  at  
7 the NADP sites. The whiskers represent the minimum and maximum values without outliers. There  
8 is a wide range of  $\delta^{15}\text{N}(\text{NO}_x)$  values within each month, with a minimum during January (-7.8~ -  
9 4.1‰) and a maximum during August (-11.4~-4.4‰). The seasonal trend shows low  $\delta^{15}\text{N}(\text{NO}_x)$   
10 during summer, with the median around -7.4‰, and high  $\delta^{15}\text{N}(\text{NO}_x)$  during winter, with the  
11 median around -6.0‰.

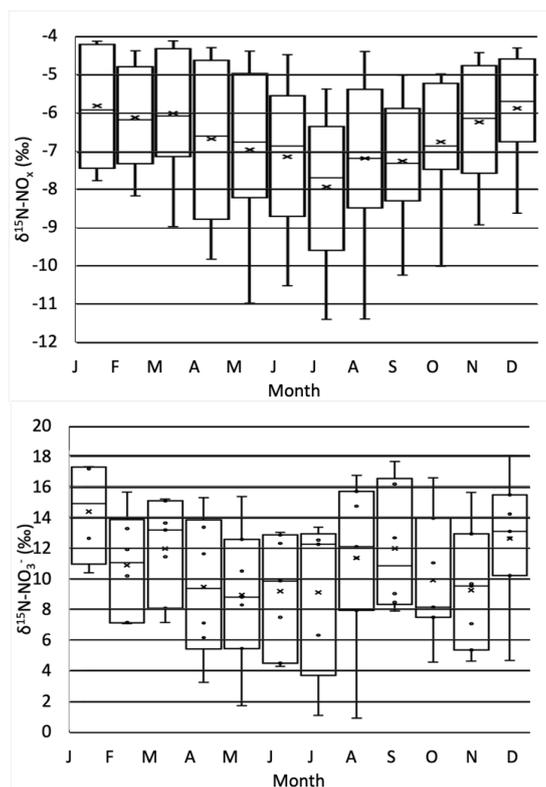


Figure 14: The CMAQ predicted  $\delta^{15}\text{N}$  value of atmospheric  $\text{NO}_x$  at NADP sites within IN, IL, OH, and KY (top) using NEI-2002 and 2002 meteorology compared to the measured  $\delta^{15}\text{N}$  of rain  $\text{NO}_3^-$  (bottom) from prior studies.

1  
2 Comparing with the CMAQ simulation, the measurements of  $\delta^{15}\text{N}$  values of  $\text{NO}_3^-$  at NADP  
3 sites from prior studies (Mase, 2010; Riha, 2013) shows the similar monthly variations and  
4 seasonal trend (Fig. 14, bottom). There is a wide range of  $\delta^{15}\text{N}(\text{NO}_3^-)$  values within each month,  
5 with a minimum during January (10.4~17.2‰) and a maximum during August (1.0~16.7‰). The  
6 seasonal trend shows low  $\delta^{15}\text{N}(\text{NO}_3^-)$  during spring, with the median around 9.3‰, and high  
7  $\delta^{15}\text{N}(\text{NO}_3^-)$  during winter, with the median around 13.0‰. The measured  $\delta^{15}\text{N}$  values of  $\text{NO}_3^-$  has  
8 the same seasonal trend as the CMAQ simulated  $\delta^{15}\text{N}$  values of  $\text{NO}_x$ . However, the measured  $\delta^{15}\text{N}$   
9 values of  $\text{NO}_3^-$  is about 17‰ higher than the CMAQ simulated  $\delta^{15}\text{N}$  values of  $\text{NO}_x$ . The difference  
10 between CMAQ simulated  $\delta^{15}\text{N}$  values of  $\text{NO}_x$  and measured  $\delta^{15}\text{N}$  values of  $\text{NO}_3^-$  is caused by the



1 following two factors: a). the mixture of isotopically lighter  $\text{NO}_x$  from the surrounding area  
2 discussed in section 3.2, and b). the net N isotope effect during the conversion of  $\text{NO}_x$  to  $\text{NO}_3^-$ ,  
3 which will be addressed in future work.

4

#### 5 4. Conclusion

6 The  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  was simulated by CMAQ, based on the emission input datasets  
7 prepared from the previous companion research (Fang & Michalski, 2020) and the meteorology  
8 input dataset simulated from WRF and MCIP.  $\delta^{15}\text{N}$  is an effective tool to track the atmospheric  
9  $\text{NO}_x$ , in term of its evolution of spatial and temporal composition, altered by atmospheric processes.  
10 The simulation indicates that the PBL height is the key driver for the mixture of anthropogenic  
11 and natural  $\text{NO}_x$  emission, which deepens the gap between  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  and  $\text{NO}_x$   
12 emission. Comparing with the measurements of  $\delta^{15}\text{N}(\text{NO}_3^-)$  from NADP sites within Indiana,  
13 Illinois, Ohio, and Kentucky, the simulated  $\delta^{15}\text{N}$  agreed well with the seasonal trend and monthly  
14 variation. The performance of CMAQ simulated  $\delta^{15}\text{N}(\text{NO}_x)$  is better than SMOKE  $\delta^{15}\text{N}(\text{NO}_x)$   
15 from the previous companion research (Fang & Michalski, 2020), due to the consideration of  
16 mixing, disperse, and transport of  $\text{NO}_x$  emission from different sources.

17 After considering the effects of  $\text{NO}_x$  emission sources and atmospheric processes, there is  
18 still an obvious gap between the simulated  $\delta^{15}\text{N}(\text{NO}_x)$  and the corresponding measurements.  
19 Therefore, before adjusting the  $\text{NO}_x$  emission inventory, the future work is to explore how  
20 tropospheric photochemistry alters  $\delta^{15}\text{N}(\text{NO}_x)$  by incorporating  $^{15}\text{N}$  into the chemical mechanism  
21 of CMAQ and comparing the simulation with the corresponding measurements. With the  
22 validation of our nitrogen isotopes incorporated CMAQ, the  $\text{NO}_x$  emission inventories could be  
23 effectively evaluated and improved.

24

25 **Data availability:** The in-detail simulation results for  $\delta^{15}\text{N}$  of atmospheric  $\text{NO}_x$  under all scenarios  
26 discussed in this paper and the CMAQ-based c-shell script for generating BCON for nested domain  
27 simulation are achieved on Zenodo.org (10.5281/zenodo.4311986). The source code for CMAQ  
28 version 5.2.1 is available at <https://github.com/USEPA/CMAQ/tree/5.2.1>. The in-detail simulation  
29 results for  $\delta^{15}\text{N}$  of  $\text{NO}_x$  emission based on 2002 and 2016 versions of National Emission Inventory  
30 and the associated python codes are achieved on Zenodo.org (10.5281/zenodo.4048992). The input  
31 datasets for WRF simulation are available at <https://www.ncei.noaa.gov/data/>.

32



1    **Author contributions:** Huan Fang and Greg Michalski were the investigator for the project and  
2    organized the tasks. Huan Fang develops the model codes, reconstruct CMAQ by incorporating  
3    <sup>15</sup>N and performed the simulation to generated  $\delta^{15}\text{N}$  values. Greg Michalski helped Huan Fang in  
4    interpreting the results. Huan Fang prepared the manuscript with contributions from all co-authors.  
5  
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