

Author's Response

We appreciate all the comments provided by the reviewers. The responses will follow the format in which both reviewers presented their comments.

NB: Lines will be referred to as “L” in Author's Response (AR), General Comments as GC, and Author's Changes as AC

Reviewer 2

GC1: A paragraph with a detailed description of how GPM works would be beneficial to readers. This would fit well within the section title “Process Modeling in GPM”.

AR1: We agree with this comment from the reviewer. Additional information on how geological processes in GPMTM operate have been included in the manuscript.

AC1: Changes that have been made in the manuscript include:

Steady Flow Process

The steady flow process in GPM model flows that change slowly over a period; e.g. rivers at normal stage, and deltas. The steady flow process best depicts sediment transport scenarios where flow velocity and channel depth do not vary abruptly. The steady flow process settings can be specified to fit a task in the steady flow pane of the “run sedimentary simulation” dialog box in Petrel software (2017.1 version and above). To attain stability in the simulator before running the full simulation (i.e. entire depositional period), it is advisable to undertake preliminary runs to ascertain the appropriateness of the source intensity and flow behaviour. For steady flow, a boundary condition must be specified at the edges of the model. In an open flow system, negative integers (i.e. values below zero) should be assigned to the edges of the hypothetical paleo-surface to allow water to enter and leave the simulation area. Further

information on the steady flow settings can be located in the GPM user manual (i.e. Guru in the Petrel software).

Unsteady Flow Process

The unsteady flow process simulates flow that are periodic, and run for a limited time; example, in turbidites where velocity of flow and depth changes abruptly over time. The unsteady flow process involves fluid elements that are affected by gravity, and by friction against the hypothetical topographic surface. A previous study on the use of unsteady flow process for stratigraphic simulation is outlined in Otoo and Hodgetts, (2019).

Diffusion Process

The diffusion process replicates sediment erosion from areas of higher slope (i.e. source location), and deposition to lower slope sections of the model area. Sediment dispersion is carried out through erosion and transportation processes that are driven by gravity. The diffusion process follows an assumption that sediments are transported downslope at a proportional rate to the topographic gradient; therefore making fine grained sediments easily transportable than coarse grained sediments. The diffusion process is controlled by two parameters; (i) diffusion coefficient, which controls the strength of the diffusion, and (ii) diffusion curve that serves as a unitless multiplier in the algorithm. The mathematical equation for the diffusion geological process is:

$\frac{\partial z}{\partial t} = k \nabla^2 z$, where z is topographic elevation, k the diffusion coefficient, t for time, and $\nabla^2 z$ the laplacian.

Sediment Accumulation

This involves the deposition of sediment using an areal source location. In the GPM™ software, sediment source can be set to a point location or considered to emanate from a

whole area. For example, where a lithology is interpreted to be uniformly distributed, the sediment accumulation process can replicate such depositional scenarios. The areal input rates (in mm/yr) for each sediment type must be specified in the settings. Specifying the areal rates for each sediment is important because the software is configured to use the value of the surface at each cell in the model and multiplies it by a value (i.e. value from a unitless curve) at each time step in the simulation to estimate the thickness of sediments accumulated or eroded from the cell. Sediment accumulation can be expressed as:

GC2: Statistical validation of why the number of modelling scenarios were chosen would be good to include. For this to be reproducible (which should be the aim), anyone that reads this should have a clear idea as to why 20 scenarios were chosen so this method can be repeated in other studies.

AR 2: A major limitation in the FSM approach is that initial boundary conditions at the time of deposition, which is required for the simulation, are unknown. In our opinion, a better means to evaluate the stratigraphic scenarios selected should be the capacity of their resultant stratigraphic-based porosity and permeability property model to match known data.

AC2: An initial simulation (labelled figure 6a) was undertaken to see if the outcome will mimic the depositional pattern observed in the seismic section (figure 3b). The 20 scenarios were derived by using different input parameters with Figure 6a as guide.

GC3: I would advise the authors to not mention any ‘future studies’ or further work. This manuscript should stand alone and showcase the modelling methods presented rather than putting a final statement about what they want/are going to do in the future.

AR 3: The suggestion will be considered, and the necessary corrections made in the manuscript.

AC 3: Further explanations on how results were achieved in GPM and integrated into the property modeling workflow in Petrel have been included in the manuscript. In addition, the mention of future studies in the manuscript have been removed.

Line Comments (LC)

L14: “where” should be used instead of “were”

AR14: The appropriate word will be used in the manuscript.

AC14: The new statement reads “Reservoir modeling techniques with the capacity to integrate forward stratigraphic simulation outputs with stochastic modeling techniques for subsurface property modeling”

L15: “accommodation space” is not a widely accepted term anymore, instead please use “accommodation”.

AR15: The appropriate word “accommodation” will be used henceforth.

AC15: “Accommodation” has been used instead of “Accommodation space” anywhere it was found in the manuscript.

L6-10: These statements do not make much sense, even with the suggested revision by the author. I would suggest something like “Typically, reservoir modelling procedures require continued property modification until a satisfactory match to known subsurface data is achieved. However, acquisition of subsurface data is costly, thus prohibitive to data collection and reservoir model conditioning.”

AR 6-10: The suggestion has been accepted and incorporated into the manuscript.

AC 6-10: Typically, reservoir modeling tasks require continued property modification until an appropriate match to known subsurface data is obtained. However, acquisition of

subsurface datasets is costly, thus restricts data collection and subsurface modeling condition; hence reducing our perspective of reservoir property variation and its impact on fluid behaviour.

L 16: This is repeated throughout the manuscript, the statement ‘most likely’ does not fit with the assertions made in a scientific manuscript. Integration of FSM’s with stochastic modelling techniques will improve reservoir characterisation because they more accurately simulate the geology than other methods.

AR 16: We agree with the reviewer’s suggestion, and will make the changes in the manuscript.

AC 16: The correction will be made to read “Reservoir modeling techniques with the capacity to integrate forward stratigraphic simulation outputs with stochastic modeling techniques for subsurface property modeling will improve reservoir heterogeneity characterization, because they more accurately produce geological realism than the other methods (Singh et al. 2013)”.

L 41: I presume the author means ‘tidal’ processes. If so, “tidal” is a more appropriate term to use in this instance.

AR 41: The appropriate changes will be made to conform with the reviewer’s comment.

AC 41: complex depositional architecture of waves, tidal and fluvial processes; suggesting that a single.

L 73: “twenty four suite of well data” should be changed to something like: “and a suite of 24 wells that comprise of..”

AR 73: We agree with the suggestion, and will make the corrections in the manuscript.

AC 73: Datasets include 3-D seismic data, and a suite of 24 wells that consist of formation pressure data, core data, and sedimentological logs.

L 76: The author states that a variety of geological features (grain size, sedimentary structures etc.) “play a significant part of reservoir petrophysics”. This is an important statement given the nature of the study, but this point should be elaborated. A sentence is all that is needed.

AR 76: Grain composition, and structure does control petrophysical attributes in a reservoir, hence some additional explanation to make the statement clearer.

AC 76: Grain size, sediment matrix and the degree of sorting generally controls the volume of voids created in a sedimentary formation, and therefore the porosity and permeability attributes.

L 101 & 106: The second sentence should start with “For example..” and Remove “space” from “accommodation space” – ‘space’ is implied in ‘accommodation’ respectively.

AR 101 & 106: The corrections will be made as suggested.

AC 101 & 106: As indicated previously, “Accommodation space” has been replaced with “Accommodation” in the manuscript.

L 108: This sentence should be broken into two, there isn’t a need for a semicolon. In fact, I would recommend that the manuscript be carefully reviewed to exclude semicolons as they create long sentences.

AR 108: The suggestion is accepted, and the relevant changes will be done.

AC 108: Sediment deposition, and its response to post-depositional sedimentary and tectonic processes are significant in the ultimate distribution of subsurface lithofacies units. To attain

stratigraphic outputs that fall within the depositional architecture interpreted from the seismic data, the input parameters were varied (see Table 2).

Lv164: I am not sure the correct figure is cited here. Should it be Fig. 4d?

AR 164: Figure 4d is the hypothetical topographic surface that was used to generate the “best fit” stratigraphic model in Figure 5d. So figure 5d as used in the manuscript is the appropriate figure.

L164: As a result, no changes will be made to this comment.

L 164-166: This statement needs a reference. Which shallow marine depositional sequence? The one presented by Folkestad and Sature (2006)? If so, please note the appropriate reference.

AR 164-166: We are agree, and will provide the references to support this statement.

AC 164-166: This is because, when compared to depositional description in studies such as Folkestad and Sature (2006); kiefert et al., (2011), it produced a stratigraphic sequence that mimics the depositional sequence in the shallow marine depositional environment under study.

L 176-178: This sentence should be condensed. Careful attention should be paid to grammatical errors and misspellings.

AR 176-178: The comment has been considered and corrections made.

AC 176-178: For example, shoreface lithofacies units were characterized using medium-to-coarse grained sediments, which accumulates at proximal distance to the sediment source. In contrast, mudstone units are associated to fine grained sediments that accumulate at distal section of the simulation domain.

L 180-183: This sentence needs to be condensed. High N/G zones are known to be the best reservoir quality zones/units. Instead focus on what those zones are from the previous work and data.

AR 180-183: In line with the reviewer's comment, appropriate changes will be made in the manuscript.

AC 180-183: In previous studies on the Sleipner Øst, and Volve field (e.g. Equinor, 2006; Kieft et al., 2011), Shoreface deposits were identified to make up the best reservoir units, whilst lagoonal deposits formed the worst reservoir units. Using this as guide, shoreface sandstone units and mudstone/shale units in the forward stratigraphic model were characterized as best and worst reservoir units respectively.

L 184: Replace "Statoil" with "Equinor"

AR 184: We agree to the suggestion, and will make the corrections.

AC 184: Anywhere "Statoil" was used in the manuscript will be replaced with "Equinor".

L 195: Please change the current statement to ' . . . extended to represent lithofacies. . . '

AR 195: The relevant modification will be done.

AC 195: The workflow (**Figure 2b**) used for subsurface property (e.g. lithofacies, and petrophysical) modeling in PetrelTM is extended to represent lithofacies, porosity, and permeability properties in the forward stratigraphic model.

L 203: Please change the current statement to 'Typically, pillars join corresponding...' as there are more words than necessary.

AR 203: We are in agreement with the reviewer's comment, and will make the appropriate changes.

AC 203: Typically, pillars join corresponding corners of every grid cell of the adjacent grid to form the foundation.

L 205: This sentence should be condensed. For example, it's not necessary to include the nomenclature of 'corner point gridding' in this context.

AR 205: We agree with the reviewer on this comment, and will be the necessary changes in the manuscript.

AC 205: The sentence with respect to the fault directions has been modified into "The prominent orientation of faults (I-direction) within the model area trends generally in a N-S and NE-SW direction, so the "I-direction" was set to the NNE-SSW direction to align the grid cells.

L 211: This sentence should be broken up into two statements and a clearer definition of layers is required.

AR 211: The comment has been taken on-board, and will be included in the manuscript.

AC 211: Vertical layering on the other hand defines the thicknesses and orientation between the layers of the model. Layers in this context describes significant changes in particle size or sediment composition in a geological formation.

L 212: What is the cell thickness? Are they constant across the model? How were they defined to control the vertical scale?

AR 212: A constant cell thickness of 1 (one) was used. This was done to attain an identical thickness as that generated from the stratigraphic simulation in all zones.

AC 212: Using the vertical layering scheme makes it possible to honour the fault framework, pillar grid and horizons that have been derived in the model.

L 215: What is meant by ‘finer’ cells? Please be more precise with the scale you are referring to. Porosity and Permeability Modelling.

AR 215: The use “finer cells” is to describe cells with smaller dimension. However, it has been modified to make it more meaningful to readers of this manuscript.

AC 215: The statement has been changed into: “ Upscaling: involves the substitution of fine grid cells with coarser grid cells. This is done to assign property values to cells in order to evaluate which discrete value suits each a selected data point. One advantage of the upscaling procedure is to make the modeling process faster.

L 223: What is the original cell size if it was compressed by 75%? This statement is unclear.

AR 223: The original cell size from our deduction is 143 m x 133 m x 84 m.

AC 223: The statement has been revised into “The original petrophysical model has a grid dimension of 108 m x 100 m x 63 m, and is compressed by 75.27% of cell size (from an approximated original cell size 143 m x 133 m x 84 m)”.

L 227 & 229: replace ‘wells to correspond’ with ‘wells that correspond’ and ‘actual well’ should be deleted and just what is in brackets should be used. ‘known data’ respectively.

AR 227 & 229: The suggestion have been incorporated into the manuscript.

AC 227 & 229: Option 1 was to assign porosity and permeability values to the synthetic lithofacies wells that correspond to known facies-associations as indicated in **Table 4**. The synthetic wells with porosity and permeability data are placed in-between known data locations to guide porosity and permeability property distribution in the model.

L 230: Where the petrophysical properties guided by any trend data? Facies information? How were the values populated in the model?

AR 230: Yes the petrophysical properties were guided with trend map that were derived from the major orientation of the stratigraphic framework. In this instance, the orientation was in a NE-SW direction.

AC 230: For option 2 the best-fit forward stratigraphic model was populated with porosity, and permeability attributes using the stratigraphic orientation captured in the seismic data (i.e. NE-SW; 240°) to control property distribution trends. Porosity and permeability were populated into the model by using the petrophysical modeling tool under property modeling process in Petrel™.

L 237: Please be more explicit as to what you mean by 80m and 120m? Is this TD of the well? MD? Or are they 80m and 120m spaced apart?

AR 237: 80 m and 120 m are the range of total depths (TD) of synthetic wells used in the study.

AC 237: Ten synthetic wells (SW), ranging between 80 m and a 120 m in total depth (TD) were positioned in the forward model to capture the vertical distribution of porosity-permeability at different sections of the stratigraphic model.

L 253: Figure 5 referenced need to have annotations which reflect the results discussed. For example, a line that indicates the MFS surface would be beneficial to readers.

AR 253: We agree totally with this comment, and have made some modifications to figure 5.

AC 253: Figure. 5 has been modified to indicate the maximum flooding surfaces (MFS) in the forward stratigraphic model. Its orientation has also been changed into landscape format (Figure 5).

L 258: A reference is required here with this statement. Please also remove ‘space’ and only just ‘accommodation’

AR 258: We have done the necessary correction on this line to include the references.

AC 258: This is consistent with real-world scenario where sediment supply match up with accommodation generated as a result of the relative constant sea level rise within a period (e.g. Muto and Steel, 2000; Neal and Abreu, 2009).

L 262-270: This contains comments that are related to spelling mistakes.

AR 262-270: The suggestions have been taken on board, and the necessary corrections will be done.

AC 262-270: The impact of the stratigraphic simulation on porosity and permeability representation in the model was evaluated by comparing its outcomes to the original porosity and permeability models of the Volve dataset using two synthetic well prefixed VP1 and VP2. The synthetic well were sampled at a 5 m intervals vertically to estimate the distribution of porosity and permeability attributes along wells. Considering that the original porosity and permeability model (**Figure 11a**) have undergone phases of history matching to enable well planning and production strategies in the Volve field, it is reasonable to assume that porosity and permeability distribution in the Volve field petrophysical model will be geologically realistic and less uncertain.

L 272: I’m not sure what is mean by ‘modal distribution’? Do the authors mean multi-modal distribution? Normal distribution? I would suggest calculated the statistical model of the original Volve porosity model and then the models of the validation wells.

AR 272: According to the petrophysical evaluation report of the Volve field by Equinor, porosity in the reservoir is between 0.17-0.30. Vertical sampling in some selected models

show more porosity values within this range (i.e. 0.17-0.30). This sentence is to illustrate how the FSM approach could generate outputs that are consistent with known data. I however, agree that more explained is required in the statement.

AC 272: The vertical distribution (**Figure 12**) of porosity in selected model realizations shows a large set of porosity values that range between 0.18 – 0.24. This output is consistent to porosity figures captured in the petrophysical evaluation of the Volve field (Equinor, 2016).

L 274-275: This statement needs to be reworded. Are the authors saying that stratigraphic inclination remains constant within the zones, or just other variogram parameters?

AR 274-275: We have made some changes to the statement as suggested by the reviewer.

AC 274-275: In view of the limitation in making variations within a simulation run in GPM™, the forward stratigraphic-based model was derived with an assumption that variogram parameters, stratigraphic inclination within zones are constant in each simulation run. In contrast, the original petrophysical model involve other measured attributes within the stratigraphic zone, hence the variations noted in Table 5b.

L277-281: This sentence is too long. Please break it into smaller, clearer sentences. Are the authors suggesting that the FSM is reasonable? Or are they suggesting that the permeability models should be conditioned to known subsurface data? This sentence should be a statement about the results of the model and what is suggested as uncertainties to consider when using these types of modelling methods.

AR 277-281: The sentence has been condensed to make it clearer to a reader.

AC 277-281: Typically, a petrophysical model like the Sleipner Øst and Volve field model will take into account other sources of data. For example, data from a special core analysis

(SCAL) will improve the reservoir petrophysics assessment. On the basis that the FSM approach did not involve these additional information from the formation, it is reasonable to suggest that the forward stratigraphic-based porosity and permeability models have been adequately conditioned to known subsurface data.

L 295-297: Sequence stratigraphy is a key component to lithofacies distribution characterisation, yes. This sentence should be condensed and reworded.

AR 295-297: Suggestion has been taken on board, and the appropriate modifications made.

AC 295-297: Indicated in previous studies, (e.g. Allen and Posamentier, 1993; Ghandour and Haredy, 2019) sequence stratigraphy is vital in the characterization of lithofacies in shallow marine settings. Aimed at replicating stratigraphic sequence formation in 3-D, the forward stratigraphic modeling approach in GPMTM provide a good framework to analyse petrophysical property variations in a reservoir.

L 302-304: These lines refer to spelling mistakes.

AR 302-304: The corrections have been made as suggested in the comment.

AC 302-304: we concede that there is a possibility to overestimate and or underestimate porosity and permeability properties as observed in some sampled intervals of the validation wells. In view of this, it is our suggestion that forward stratigraphic simulation outputs should be applied as additional data to understand sediment distribution patterns, and associated vertical and horizontal petrophysical trends in the depositional environment than using its outputs as an absolute conditioning data in subsurface property modeling.

L 314-315: This sentence is too long. Please split into two statements for clarity.

AR 314-315: The sentence has been condensed as suggested by the reviewer.

AC 314-315: In reality, sediment deposition into a geological basin is also controlled by mechanical and geochemical processes that tend to modify a formations petrophysical attributes (Warrlich et al. 2010). Therefore, using different geological processes and initial conditions to generate depositional scenarios, will help to produce a best fits stratigraphic framework of the reservoir under study.

Reviewer 1

GC1: A big general concern I have regarding the manuscript is that all results hinge on the realistic prediction of the sediment deposition by the stratigraphic model GPM, but GPM is in this manuscript described and treated as a kind of "black box". As a reader, without knowledge of how GPM works internally, there is no way to check, know or estimate, how, why, or if the given input parameters will yield the results presented in the manuscript.

AR 1: We agree with the reviewer that the GPM software (at least the versions I used; 2017.1 to 2019.1 versions) acts as a "black box". However, the software tries to replicate a real world sedimentary process. For example, increasing sea level, and subsidence rate in GPM corresponds with an increase in accommodation. Similarly, high land-ward elevation, and erosion increases sediment supply into the basin.

AC 1: The focus of this study is to produce a depositional sequence that mimic the pattern in the seismic section in Figure 3b. Throughout the manuscript, further information will be provided in relevant sections to bring clarity on how the GPM software works.

GC 2: I am missing at least some basic equations or general explanations how the geological processes (Sediment Diffusion, Tectonics, Steady Flow, Unsteady Flow and Waves) act in GPM on the given input parameters. I, as someone not having worked with GPM before, have no idea and no possibility to understand how exactly the simulation results result from the little descriptions in the text and the values given in Table 2. What does GPM assume how the geological processes (Sediment Diffusion, Tectonics, Steady Flow, Unsteady Flow and Waves) act on the sediments? Another issue I have with the manuscript is that it is not made clear enough how the 20 scenarios of the stratigraphic simulations are connected with

the 50 realizations. What are the realizations? How are they generated? For which scenario(s) are they applied? How and why are parameters of the scenarios and realizations chosen?

AR 2: This comment is very similar to the observations of reviewer 2. In addition, 20 scenarios are because of the uncertainty associated with the input parameters used to for the simulation. Different inputs were used to obtain a most representative stratigraphic framework. The 50 realizations are generated in the property modeling stage (porosity and permeability) in Petrel software, where synthetic wells data from a simulation (in this instance scenario 4) are used in a geostatistical algorithm (i.e. sequential Gaussian simulation) to generate a range of property representations. This will allow us to compare which outcome(s) match the original Volve field model from Equinor.

AC 2: We totally agree with the comment and the modification in the manuscript you follow the same style as the one presented in AC 1 under reviewer 2.

GC 3: In lines 86-87 the authors state "Simulations were constrained to twenty scenarios because the desired stratigraphic sequence and associated sediment patterns were achieved at the fourth simulation." I miss the discussion or details on how the authors determine what a desired stratigraphic sequence looks like and why they continued for 16 more scenarios, when Scenario 4 was already giving the desired stratigraphic sequence?

AR 3: A desired stratigraphic pattern in this contribution is one that exhibits similarity to the depositional sequence observed in the seismic section shown in Figure 3b. Additional 16 scenarios were generated as an attempt to enhance the results in scenario 4.

AC 3: Further explanations will be made in relevant sections to make the statement clearer to readers of this manuscript.

GC 4: A final general comment is that the authors claim that the presented approach reduces the uncertainty in the distribution of petrophysical properties such as porosity and permeability. In the conclusion, the authors discuss that even with their approach, uncertainty in the distribution of petrophysical properties will remain. They might increase the impact of their paper and prove their claim by comparing their resulting distribution with ones that are generated by other, more classical methods. Although I acknowledge that this might be a too big of a topic to include in this manuscript, I would find it good to at least mention the possible interest of such a comparison in the conclusion.

AR 4: In a related work, which compares the forward stratigraphic-based modeling approach to a classical technique (e.g. pixel based modeling) is being worked on. Notwithstanding that, this suggestion will be highlighted in the concluding part of this manuscript.

Line Comments

L 6-8: Something missing in this "Typically, reservoir modeling requires property-modifying coefficients in the form values to achieve a good match to known subsurface well data."

AR 6-8: The statement is not conclusive and not clear, so the appropriate changes will be made.

AC 6-8: Typically, reservoir modeling require continued property modification until an appropriate match to known subsurface data is obtained.

L 25-27: Something missing in this "but the method tends to confine reservoir property models to known data and rarely realize geological realism to capture sedimentary that have led to reservoir formation" → sedimentary processes??"

AR 25-27: Again, we agree the statement is not conclusive. The corrections will be done as suggested.

AC 25-27: This statement will be modified to read "but the geostatistical-based method tends to confine reservoir property models to known data and rarely realize geological realism to capture sedimentary that have led to reservoir formation".

L 39: Something missing in this "The sedimentary system, Hugin formation makes up the main reservoir interval in the Volve field."

AR 39: We have taken the suggestion on board, and modification will be made to reflect the comment.

AC 39: The reservoir interval under study is located within the Hugin formation, which studies by Varadi et al. (1998); Kieft et al. (2011), have shown to be a complex depositional architecture of waves, tidal and fluvial processes; suggesting that a single depositional model will not be adequate to produce a realistic lithofacies distributions model.

L 69-70: Something missing in this "but the thickness have not been completely penetrated (Folkestad & Satur, 2006).

AR 69-70: The relevant changes will be made in the manuscript.

AC 69-70: This statement will be revised into “but the total thickness of code F lithofacies is not known (Folkestad & Satur, 2006).”

L 86-87: The desired stratigraphic sequence and associated sediment patterns were achieved" How did you determine this? What was the criterion for this decision? and then, why did you add another 16 scenarios, if the 4th was already showing "the desired stratigraphic sequence"? The scenarios are never discussed in detail and hidden away in Table 2 and only hinted at in some short statements e.g. lines 126-127 "To mitigate this uncertainty, 5 paleo topographic surfaces were generated stochastically" or lines 148-149 "The simulation parameters applied (Table 2) were generated randomly using the initial run.

AR 86-87: The main criteria for evaluating the realistic nature of a stratigraphic model was to compare it to the depositional sequence observed in the seismic section in Figure 3b, and/or interpreted through well correlation.

AC 86-87: The changes made in these lines was to include the Author’s response (AR) above.

L 99-114: As said in the general comments, I miss some detail on how the mentioned processes are implemented in GPM, how are they parameterized etc...

AR 99-114: This is related to GC2, so the explanation provided will suffice for this comment.

AC 99-114: The changes that have been made include details on the parameterization of the geological processes used in simulation.

L 128: TPr is not defined. I assume it is the " paleo topographic surface" or something similar from the context.

AR 128: Yes, TPr is the paleo-topographic surface of the model area.

AC 128: To mitigate this uncertainty, 5 paleo topographic surfaces (TPr) were generated by adding or subtracting elevations from the inferred paleo topographic surface or base topography (see Figure 4g).

L 133-136: The sediment entry point for this task was placed in the north-eastern section of the hypothetical paleo-topography. Since the exact sediment entry point is uncertain, multiple entry points were placed at 4 m radius around the primary location in (Figure 3c), in order to capture possible sediment source locations." Compared to the scale shown in Figure 3c or the area given in line 221 (~18km²), a 4m radius seems to me just as the same location, as the modelled area seems kilometres wide. Or do you mean 4km? Could an sediment entry point actually be as narrow as 4m within such a relatively flat looking domain as shown in Figure 3c.

AR 133-136: The distance should be 4 km and not 4m as stated in the manuscript. The correction will be made in the manuscript.

AC 133-136: Based on regional well correlations in previous studies (e.g. Kieft et al. 2011), and seismic interpretation of the basin structure, the sediment entry point was placed in the north-eastern section of the hypothetical paleo-topography. Since the exact sediment entry point is not known, multiple entry points were placed at 4 km radius around the primary location in (Figure 3c), in order to capture possible sediment source locations. The source position is characterised by positive integers (i.e. values greater than zero) to enable fluid flow to other parts of the simulation surface.

L 139-140: What was the assumed sea level after 20000 years? Only the average sea levels are given later in Table 2

AR 139-140: The sea level curve used in the simulation followed the Haq global sea level curve generator as well as the Exxon global sea level curve generator formats. The sea level for year 20,000 was assumed to be 45 m, and decreased to 15 m by year zero. The sea level was not kept constant as it is a curve that covered a period of geological period (see figure 1). Averages were used in the manuscript to provide an insight into the mean sea level that was in the simulation scenarios.

AC 139-140: To attain stability in the simulator, we assumed a sea level that range between 15 m to 45 m; averaging 30 m for short simulation runs, e.g. 5000 to 20000 years. The sea level was varied with increasing duration of the simulation (illustrated in Table 2). The peak sea-level in the simulation represents the maximum flooding surface (Figure 5d), and therefore the inferred sequence boundary in the geological process model.

L 148-149: and following: "The simulation parameters applied (Table 2) were generated randomly" on what basis were they created? Was the simulation always constant with no changes? I did not find any boundary conditions, so how much sediment enters the study area? Is this constant over time? The following lines e.g. "A sudden change in subsidence rate tends" suggest that the (boundary) conditions changed over time.

AR 148-149: With scenario 1 (Figure 6a) beginning to show resemblance to the target output (i.e. the depositional pattern observed in seismic section; Figure 3b), we generated input figures that were higher and lesser than those used in generating scenario 1. Example, based on a diffusion coefficient of 8 m²/a that was used in scenario 1, diffusion coefficients +/- 5 of 8 were generated with the aim to improve the development in scenario 1. Since the initial

conditions (boundary conditions) at the time of deposition are unknown, an attempt was made to apply input parameters that will produce a comparable stratigraphic pattern to what is observed in the seismic section (Figure 3b). Aside the initial topography that was kept constant in a simulation run, other input parameters such as diffusion, wave event, steady/unsteady flow, tectonics use curve functions to provide variations within the simulated period.

AC 148-149: Modifications made in the manuscript is the same as the author's response (AR) above.

L 157: "Shifting the source point to the mid-section of the topography" to where exactly? can you show that in Figure 3? Isn't the sediment entry point shown in figure 3 already somewhat in the mid-section, at least when looking at figure 3? And previously, you wrote that you only look into changes within a 4m radius of the sediment entry point, so I do not understand how it can have such a big influence, see lines 133-136.

AR 157: The "mid-point" used in the manuscript is the middle section of the entire topography (i.e. basin-ward, close to the basin slope; see modified image below; labelled Figure 2). The point was made to show that the location of the sediment source in the simulation will have a huge impact on the resultant stratigraphic architecture.

AC 157: Shifting the source point to the mid-section of the topography (i.e. the mid-point of the topography in a basin-ward direction) resulted in the accumulation of distal elements that are identical to turbidite lobe systems.

L 176-178: "shoreface lithofacies units were characterized using medium-to-coarse grained sediments to that are proximal sediment source, while mudstone units are constrained to the distal parts of the stratigraphic model, where fine grained sediments accumulate at the end of

the simulation." → coarse grained sediments that are proximal to the sediment source? "at the end of the simulation." Do you mean at the distal end of the simulation domain or towards the end of the simulated time? Time or space is not clear from the wording.

AR 176-178: Here, we mean the distal end of the simulation domain at the end of each run in the GPM simulator. The appropriate modification will be done in the manuscript to make the point clearer.

AC 176-178: For example, shoreface lithofacies units were characterized using medium-to-coarse grained sediments, which accumulate at proximal distance to the sediment source. In contrast, mudstone units were restricted to fine grained sediments that accumulate at distal section of the simulation domain.

L 179-180: "attributes, which is" → "attributes, which are"

AR 179-180: These are Wireline-log attributes such as gamma ray, neutron porosity, sonic, and density logs outlined in Table 1 in the supplement.

AC 179-180: Using published studies by Kieft et al., (2011), porosity and permeability variations in the stratigraphic model were estimated from wireline-log attributes such as gamma ray, neutron, sonic, and density logs outlined in Table 1.

L 186: x not defined.

AR 186: "x" as used here is the multiplication symbol.

AC 186: $\varnothing_{er} = \varnothing_D + \alpha \cdot (NPHI - \varnothing_D) + \beta$; where \varnothing_{er} is the estimated porosity range, \varnothing_D is density porosity, α and β are regression constants; ranging between -0.02 – 0.01 and 0.28 – 0.4 respectively, $NPHI$ is neutron porosity.

Line 197-218: inconsistent numbering

AR 197-218: (i) and (ii) were used to show that the pillar gridding process, horizon, zoning and layering processes are all part of the structural modeling process. The numbering will be modified into a 1 to 4.

AC 197-218:

1. Structure modelling: identified faults within the study area are modelled together with interpreted surfaces from seismic and well data to generate the main structural framework, within which the entire property model will be built. The procedures involve modification of fault pillars and connecting fault bodies to one another to attain the kind of fault framework interpreted from seismic and core data.
- (2) Pillar gridding: a “grid skeleton” that is made up of a top, middle and base architectures. Typically, pillars join corresponding corners of every grid cell of the adjacent grid to form the foundation for each cell within the model. The prominent orientation of faults (i.e. I-direction) within the model area generally trends in a N-S and NE-SW direction, so the “I-direction” was set to the NNE-SSW direction to capture the structural description.

L 223: "and compressed by 75.27% of cell size" the verb is missing → "and is compressed by 75.27% of the cell size"?

AR 223: The sentence shall be corrected.

AC 223: The original petrophysical model has a grid dimension of 108 m x 100 m x 63 m, and is compressed by 75.27% of cell size (from an approximated cell size 143 m x 133 m x 84 m). To achieve a comparable model resolution as the original porosity and permeability

model, the forward stratigraphic output with initial resolution of 90 m x 78 m x 45 m was upscaled to a cell size of 107 m x 99 m x 63 m.

L 237: What are the length measures? Well lengths, distances, ...?

AR 237: This statement will be corrected to make its meaning clearer.

AC 237: Ten synthetic wells (SW), ranging between 80 m and a 120 m in total depth (TD) were positioned in the forward model to capture the vertical distribution of porosity-permeability at different sections of the stratigraphic model. The average distance between these wells as shown in Figure 9c is about 0.9 km apart, with a maximum and minimum of 1.3 km and 0.65 km respectively.

L 243: "populated" -> populate + How can wells be upscaled to the original structural model? Upscaling usually refers to representing something at a larger scale, not to extrapolate from lower dimensional objects (wells are practically 1D) to higher dimensions (the 3D structural model). I am confused here, but my guess is that the 1D to 3D extrapolation is meant here with upscaling. Please clarify. After rethinking, I do not even understand the purpose of the 10 synthetic wells, why do you use them? As I understand it, you should have from the previous steps already the full 3D structural stratigraphic information, so why throw away all that, keep only 10 locations and then reconstruct again everything? Couldn't you just directly populate the stratigraphic 3D domain?

AR 243: The synthetic wells derived from the stratigraphic model is to provide an additional well data for use in a traditional modeling workflow as was the case in the building of

original Volve model. Using the same structural model was to attain a comparable framework for evaluating the modeling outputs. Upscaling the synthetic well data is a standard procedure to “transform” the data from 1-D into a 3-D framework to build the property model.

AC 243: The variogram model (**Figure 10**), of dominant lithofacies units in the formation served as a guide in the estimation of variogram parameters from the forward model.

L 249-250: "Out of fifty model realizations, six realizations that showed some similarity to the original petrophysical model are presented" How did you generate the 50 realizations exactly? How did you quantify the similarity? For which scenario did you do the 50 realizations? All 20? Only scenario 4? Could you at some point specify this, so for what scenarios did you do the model realizations? And I assume you mean the "Porosity and Permeability model", can you confirm?

AR 249-250: The selection of six realizations was based on visual and statistical comparison of zones in the original Volve field model, and the stratigraphic-based porosity/permeability models. The statistical approach involved the comparison of summary statistics from the original Volve model, and the model realizations generated in the Petrel software. The visual comparison on the other hand looks at how geological realistic the output is, and if it conforms with our conceptual idea of the Volve field model.

AC 249-250: Out of fifty model realizations, six realizations that showed some similarity to the original petrophysical model are presented (**Figure 11**). This was achieved through visual and statistical comparison of zones in the original Volve field model, and the stratigraphic-based porosity/permeability models. The statistical approach involved the comparison of

summary statistics from the original Volve model, and the model realizations generated in the Petrel software. The visual comparison on the other hand looks at how geological realistic the output is, and if it conforms with our conceptual idea of the Volve field model.

L 277-278: Did you do any of that what you write is "typically" done?

AR 277-278: No we didn't do that. The explanation to this is that a property model that has been used for production purposes would have gone through different phases of history matching, hence its adoption as a reasonable base model. The aim is to ascertain the practicability of using the forward stratigraphic modeling technique to predict property variation in a hydrocarbon reservoir.

AC 277-278: Typically, a petrophysical model like the Sleipner Øst and Volve field model will take into account other sources of data. For example, data from a special core analysis (SCAL) will improve the reservoir petrophysics assessment. Considering that the FSM approach did not involve these additional information from the formation, it is reasonable to suggest that the forward stratigraphic-based porosity and permeability models have been adequately conditioned to known subsurface data.

L 291: "multiple simulation scenarios" The 20 (GPM?) simulation scenarios defined in Table 2? How do they link to the poro-perm model realizations? See comment on lines 249-250.

AR 291: The 20 simulation scenarios generated are related to the depositional models (stratigraphic models). Out of the 20 scenarios, scenario 4 was adopted and populated with porosity and permeability attributes. So out of the 20 stratigraphic modeling scenarios only scenario 4 has a direct relationship to the 50 realizations produced in the property model.

AC 291: Since the initial conditions of this basin is uncertain, multiple simulation scenarios were carried out to account for the range of bathymetries that may have influenced sediment transportation to form the present day Hugin formation. The simulation produced well defined clinoforms and sequence boundaries that depict the pattern observed in the seismic data. Clinoforms in this context, are sloping depositional surfaces in a stratigraphic architecture (Patruno & Hansen, 2018).

L 298: "A porosity-permeability model that match the original petrophysical model was produced" → A porosity-permeability model matching the original petrophysical model was produced.

AR 298: We agree to the suggestion from the reviewer, and will make corrections to that effect.

AC 298: A porosity-permeability model matching the original petrophysical model was produced using synthetic porosity and permeability logs from the forward stratigraphic model as input datasets in the sequential Gaussian simulation algorithm.

L 340: "will improve property prediction away from data" Away from data sounds weird to me, what do you mean with that exactly? Extrapolation away from points (wells) where there is data (well logs)?

AR 340: More conditioning data (well data) will enhances the chance of attaining realistic distributions in the model area. So with the forward stratigraphic-based property model providing a realistic stratigraphic framework, synthetic wells can be obtained to control property modeling of the reservoir. In addition the term "data" used in the manuscript refers to well logs.

AC 340: The good match obtained from validation wells in the original and stratigraphic-based petrophysical model, leads us to the suggestion that an integration of variogram parameters from well data and forward stratigraphic simulation outputs will improve property

prediction in inter-well zones. This suggestion is supported by the idea that more conditioning data (well data) will increase the chance of producing realistic property distribution in the model area.

Line 355-358: How can you guarantee that the artificial neural network approach will not have similar biases, which only are better hidden as they are less understood? How do you provide training data without cognitive or sampling biases to ensure that the artificial neural network will not train to reproduce those biases?

AR 355-358: In our view, the calculator approach used in estimating the lithofacies proportions in the stratigraphic model were constrained to the extent to which we assume such distributions should go. Meanwhile, with an unsupervised machine learning via neural network, we can attain many outcomes that are not restricted by our cognitive biases. The neural network can be defined with varying components (e.g. weights) to attain different outcomes, from which a best fit vertical profile that is comparable to real well log can be adopted.

List of Figures

Generally, comments on figures had to do with its clarity and caption. The appropriate corrections have been done, such that the orientation of some of the figures have been changed into landscape to make them clearer.

Also, Figure 12 have been divided into two (i.e. Figure 12a, and Figure 12b), in order to make it clearer and readable.

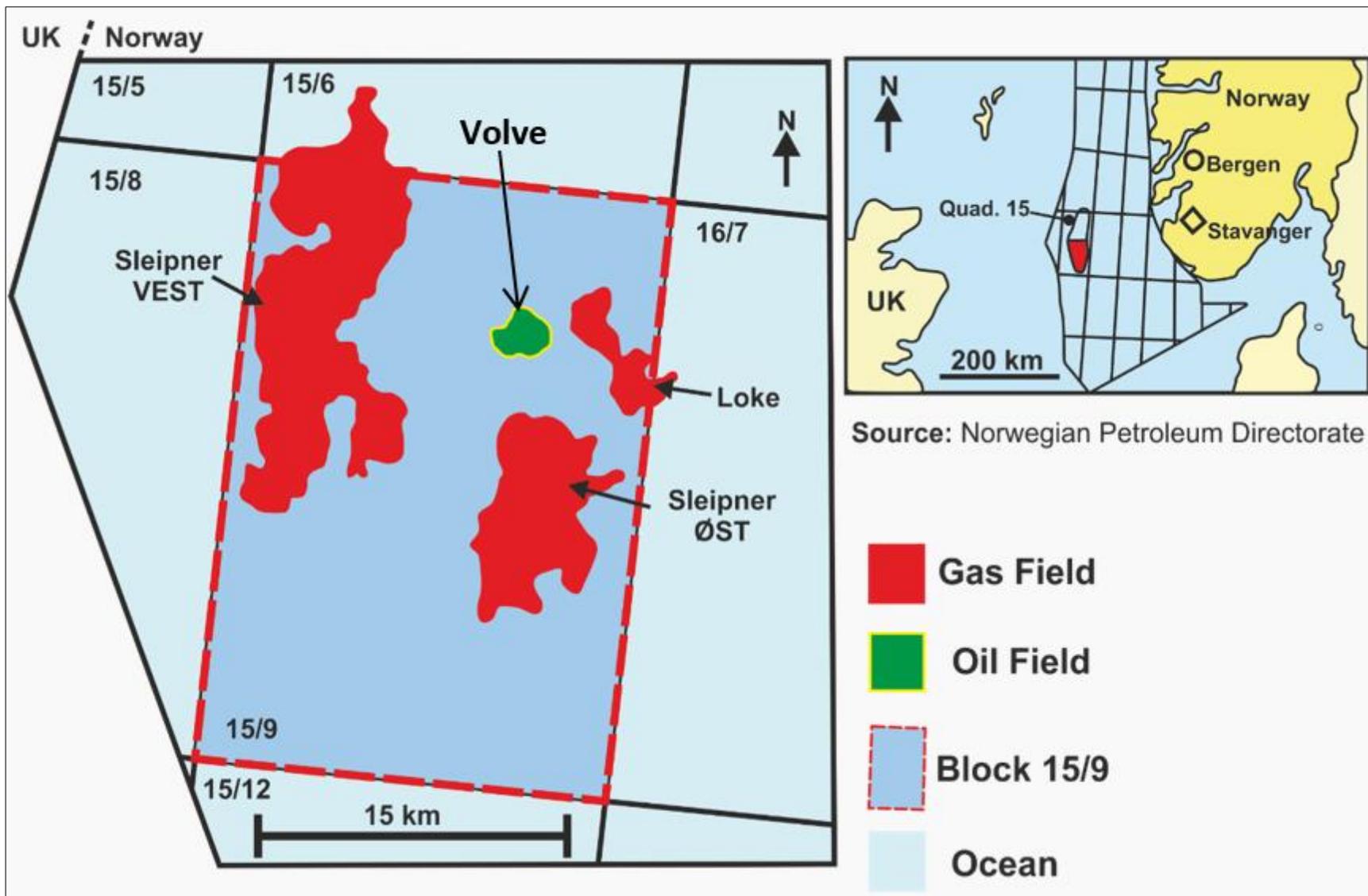


Figure 1

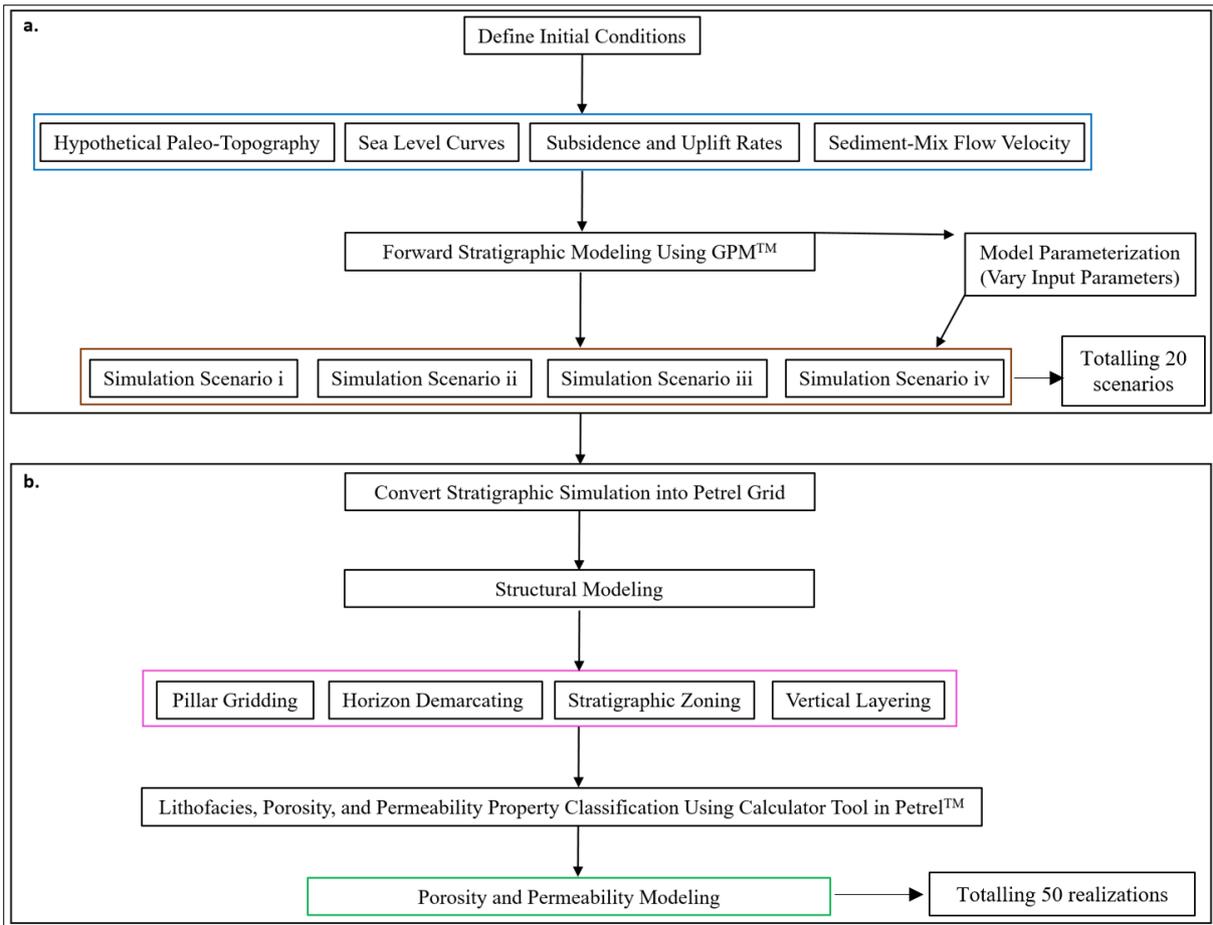


Figure 2

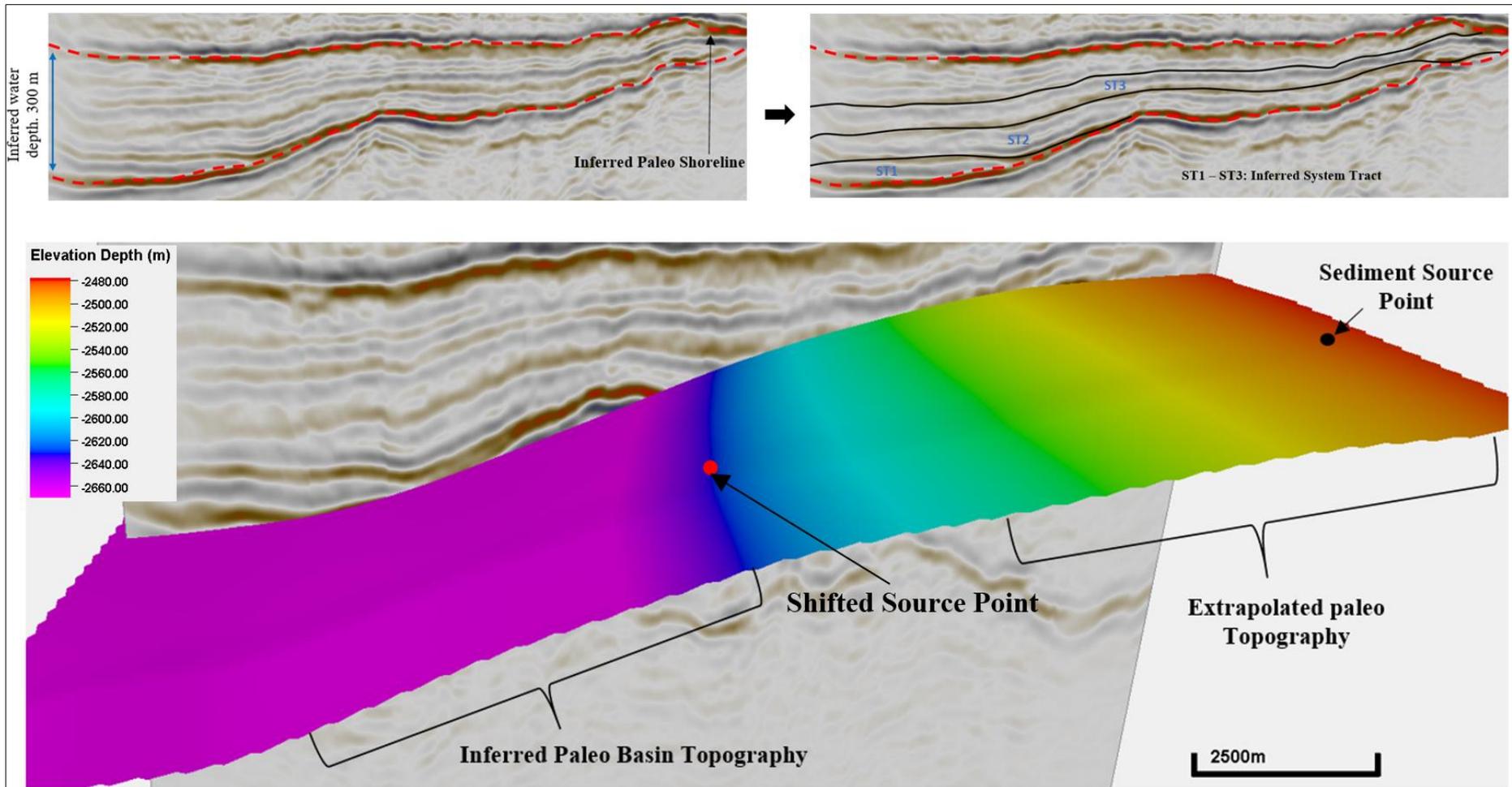


Figure 3

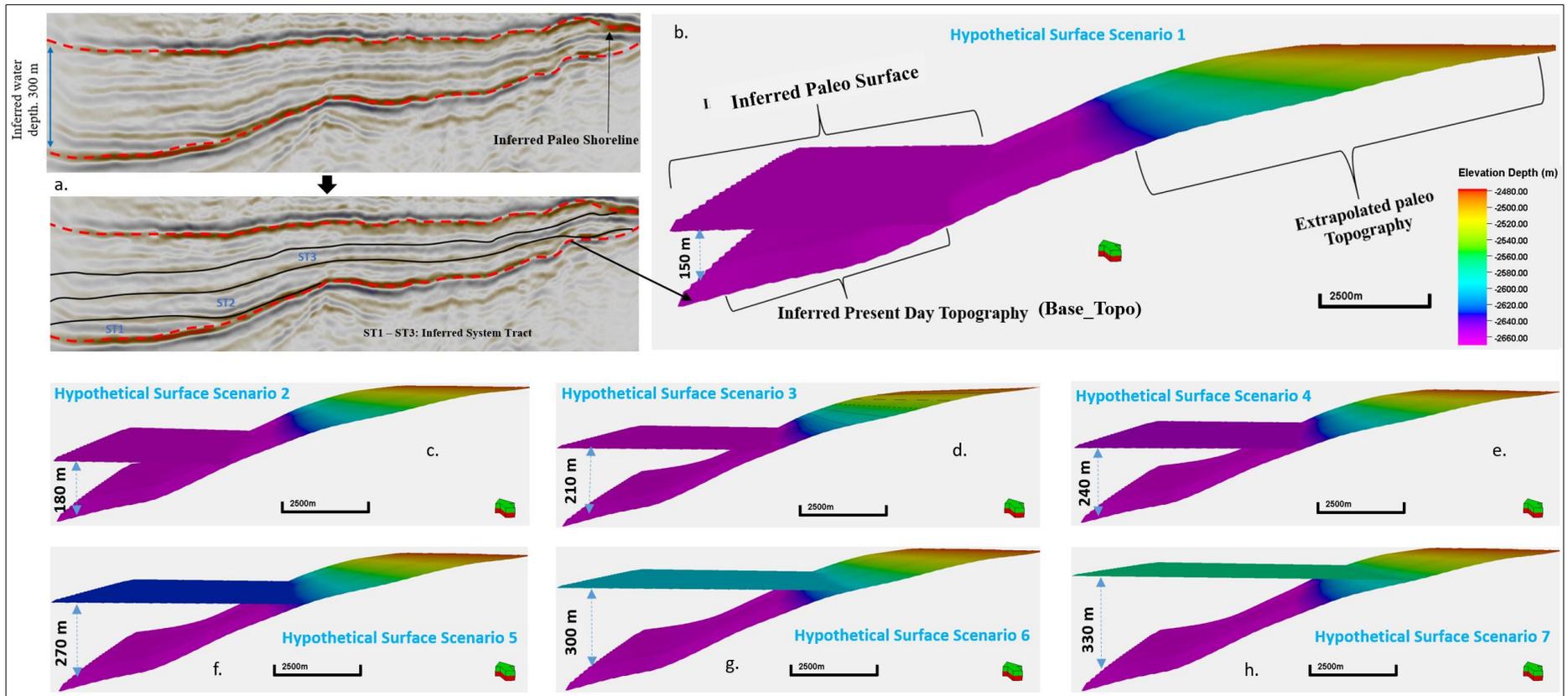


Figure 4

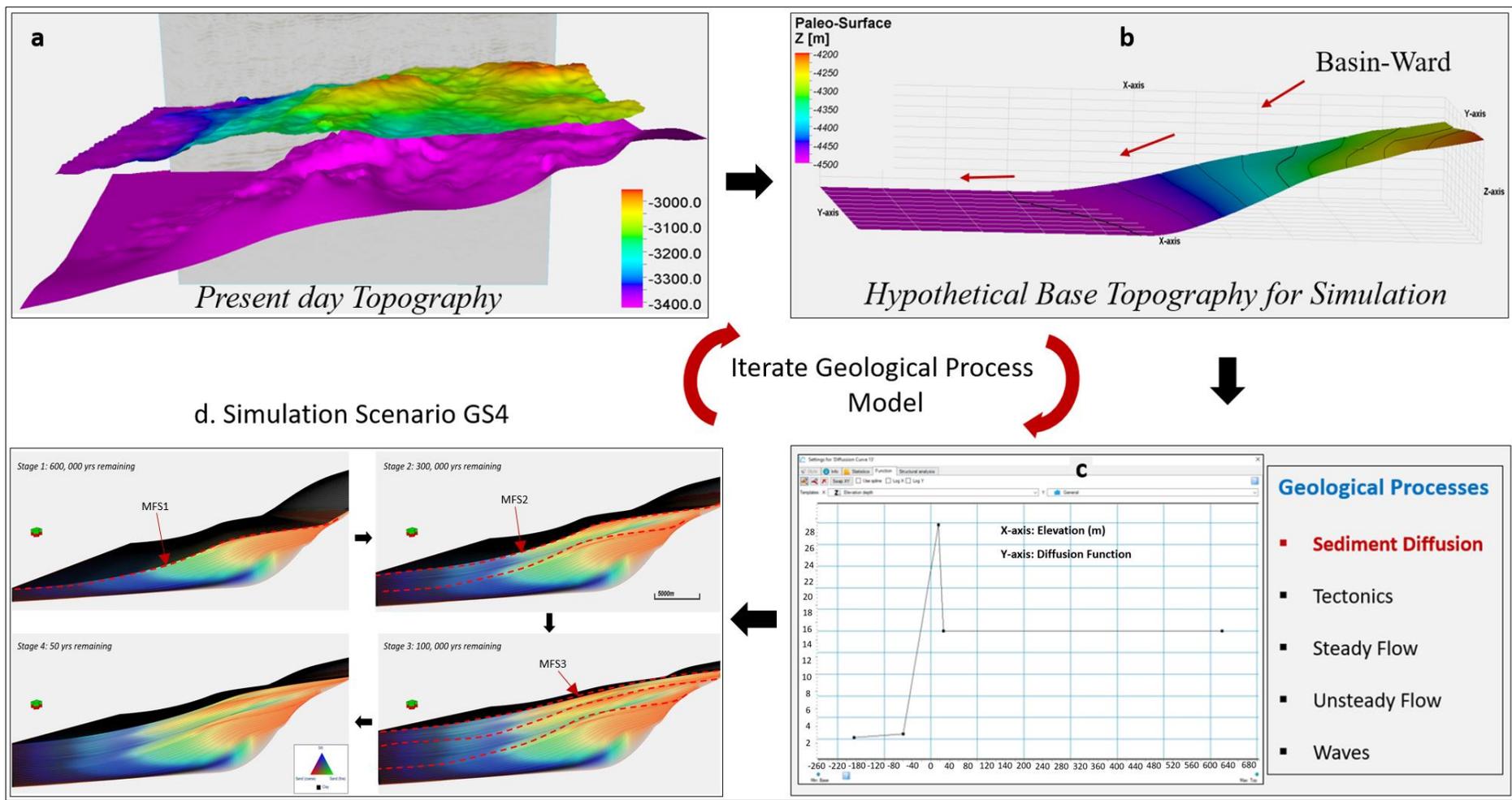


Figure 5

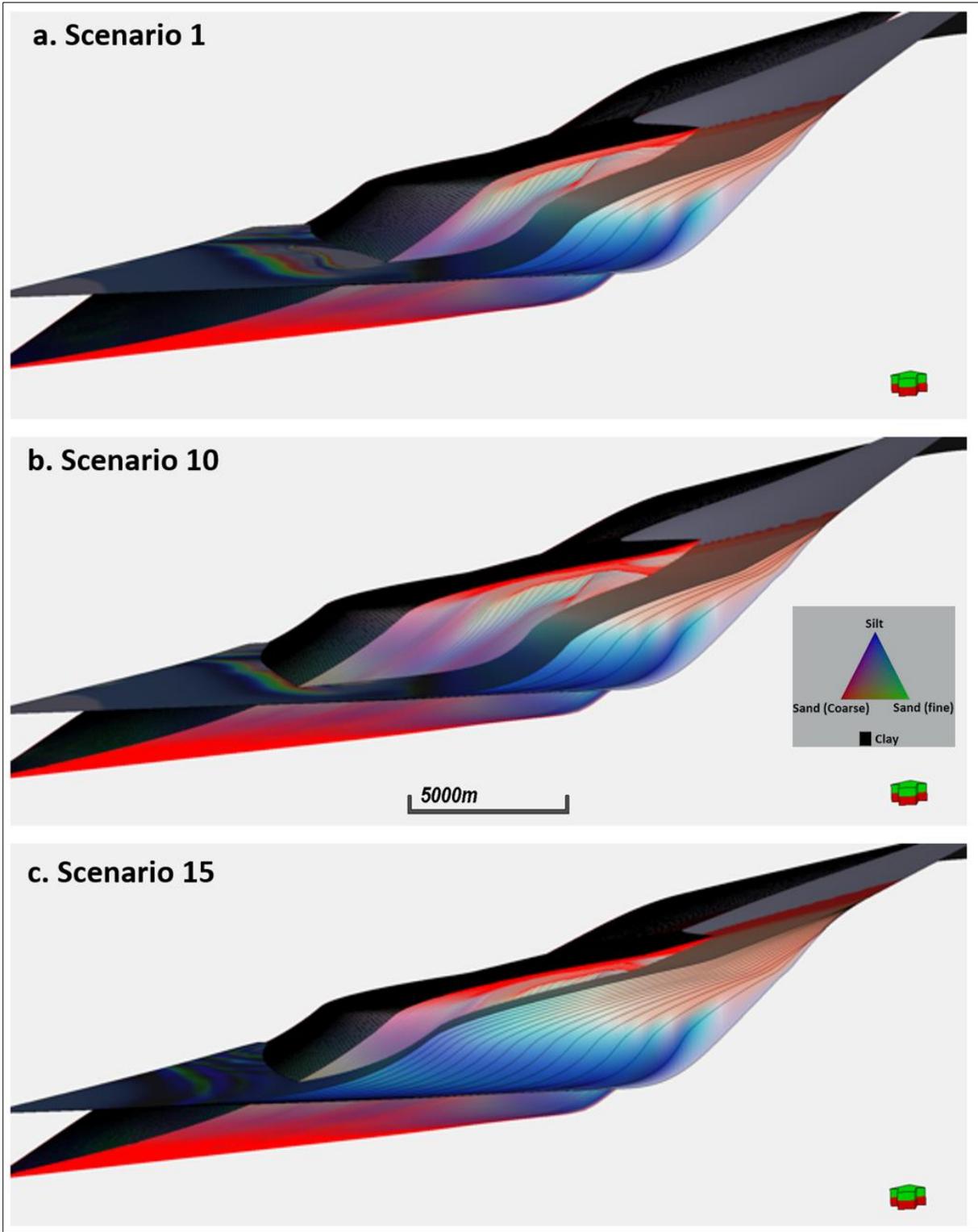


Figure 6

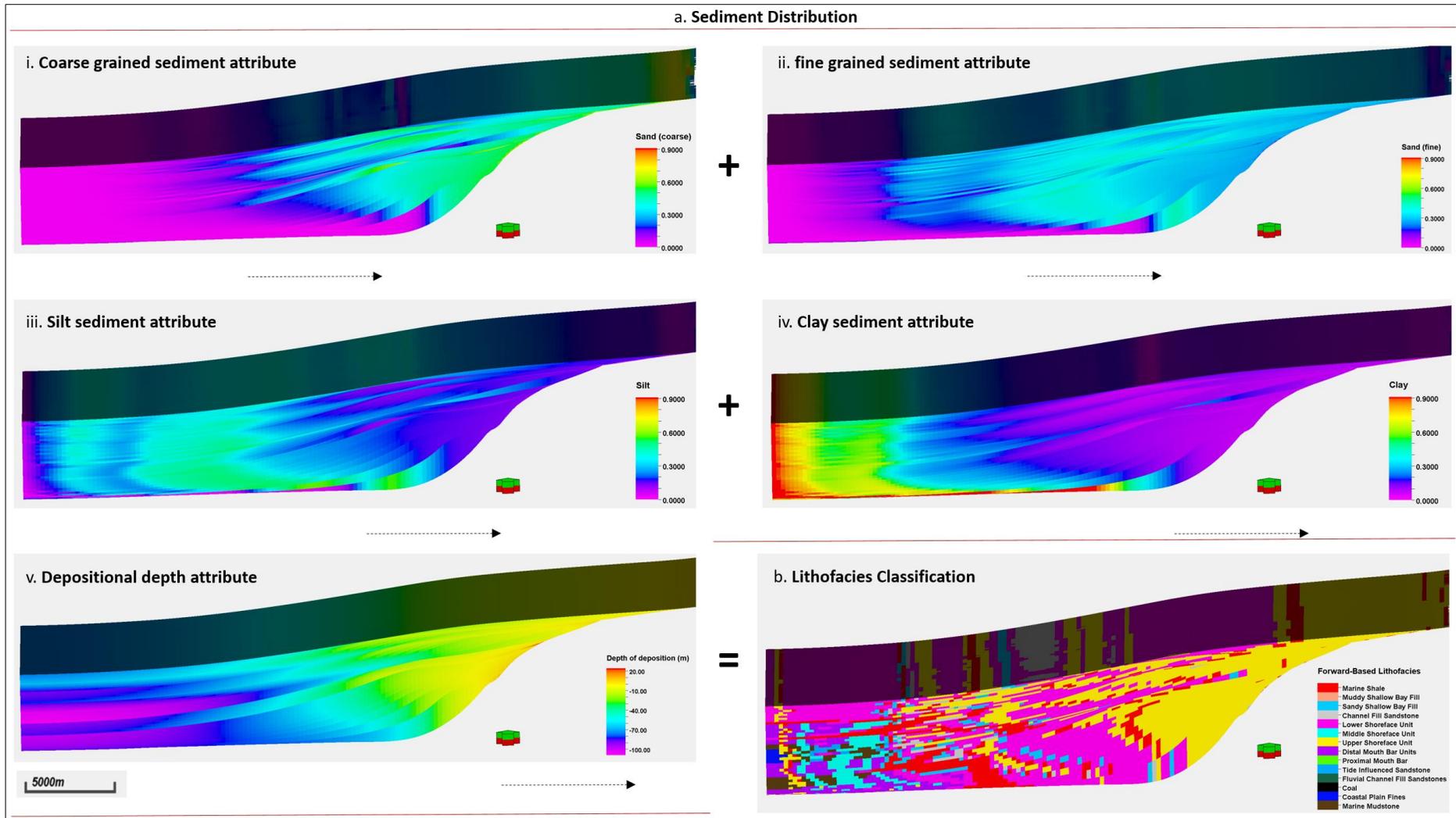


Figure 7

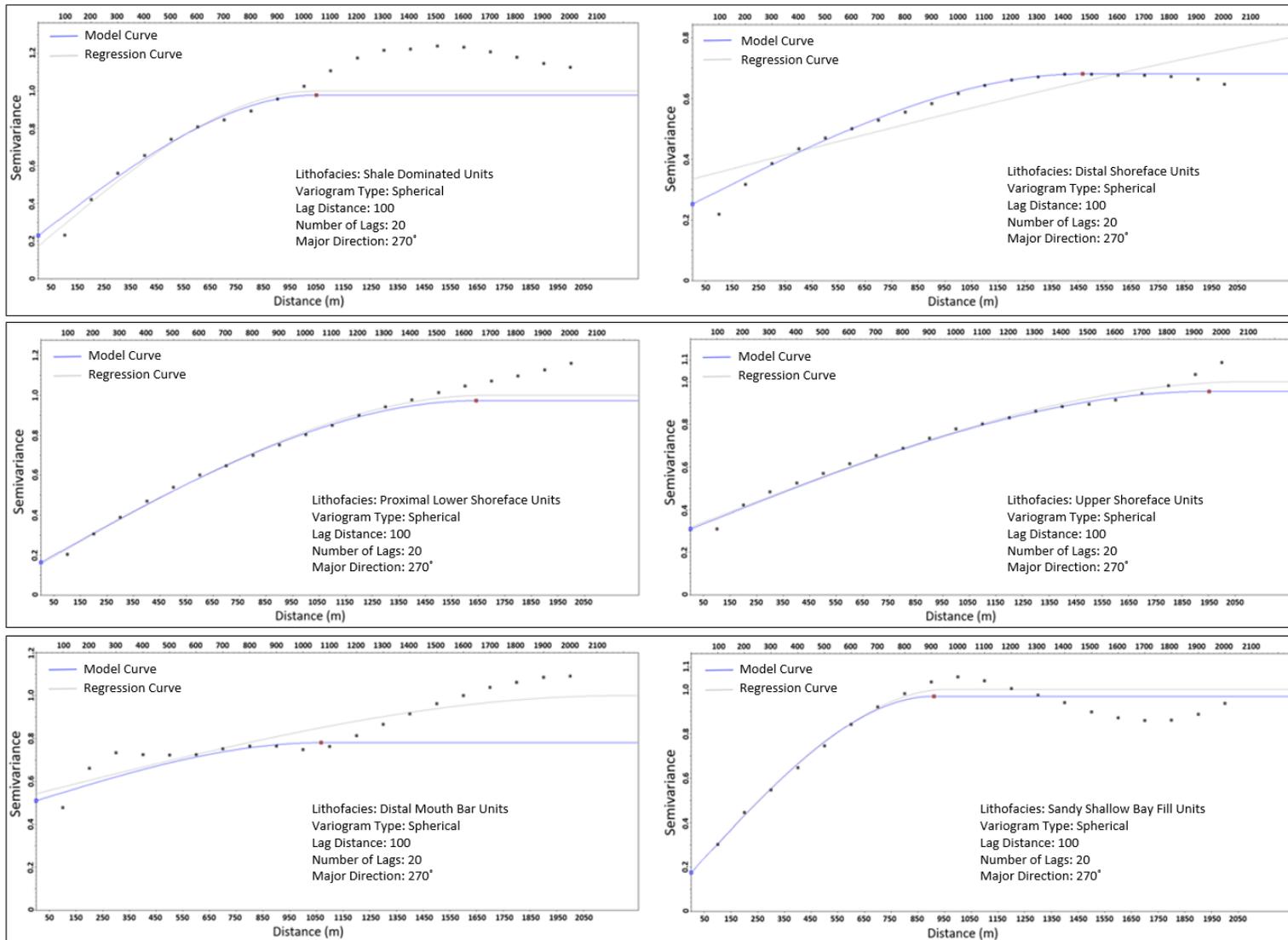


Figure 10

a. Validation Well 1

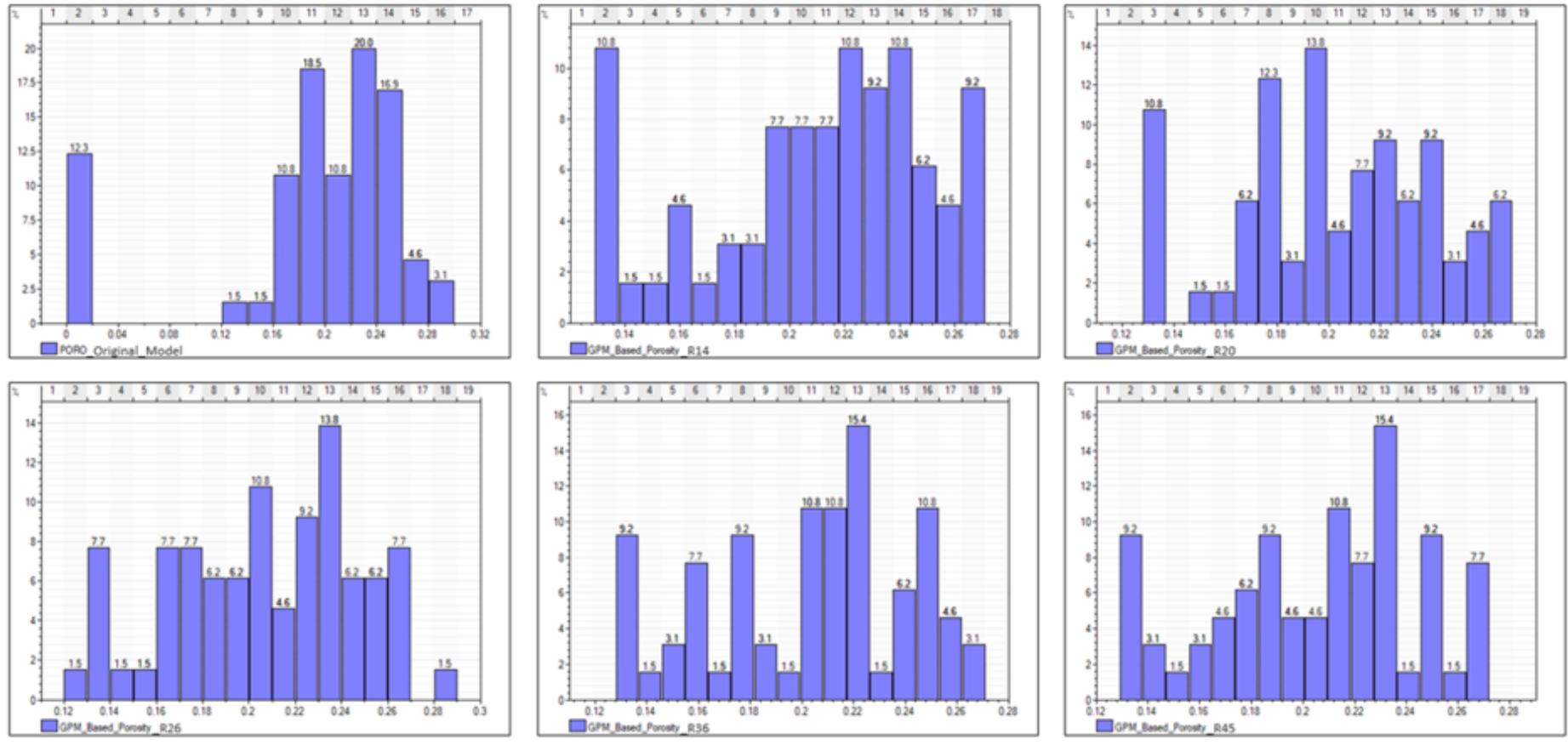


Figure 12a

b. Validation Well 2

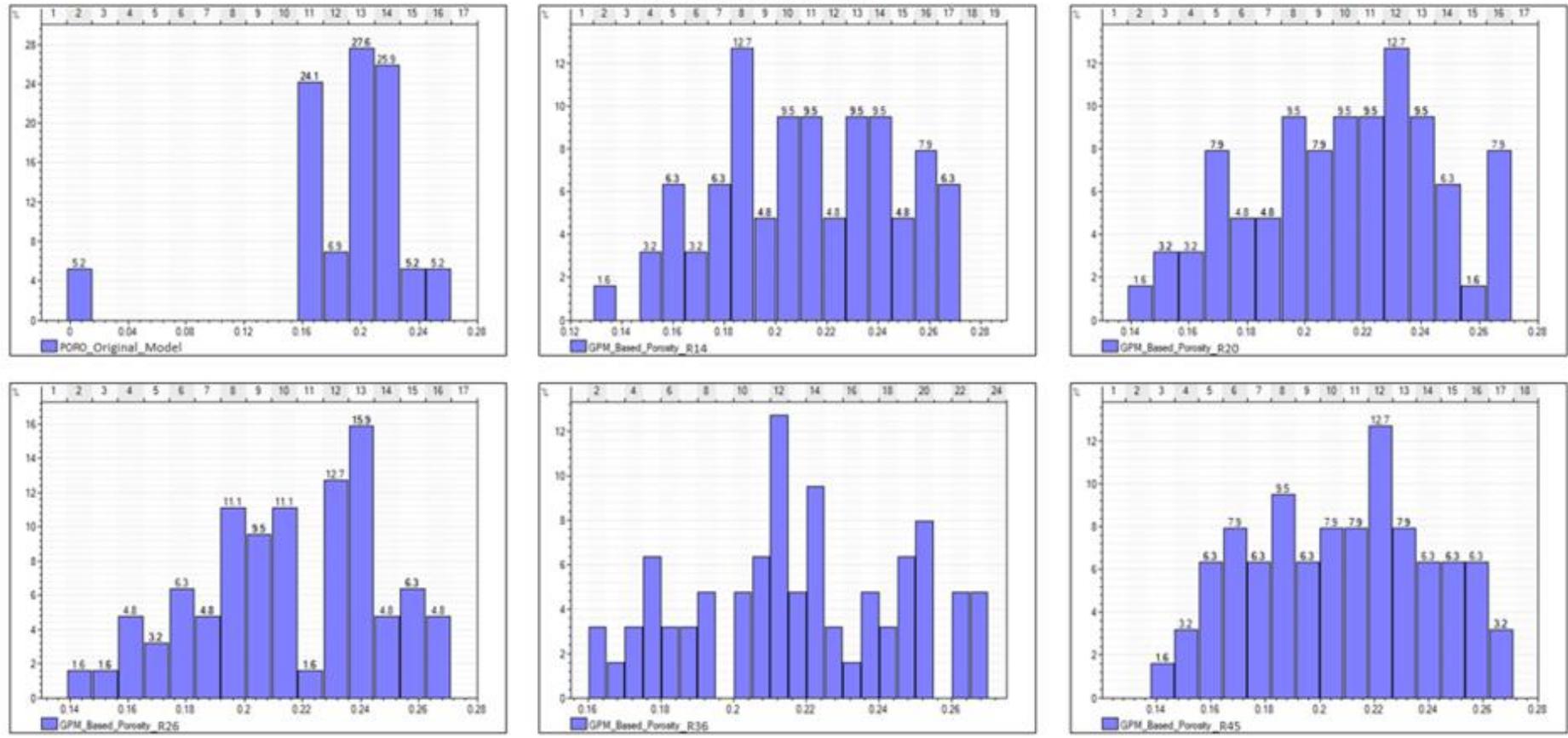


Figure 12b

Table 1

Code	Facies	Description	Thickness (t); extent (l)	Wireline-log Attribute	Interpretation
A	A1	Parallel-laminated mudstone with occasional siltstone inputs. Monospecific pattern of disorder bivalves parallel to bedding.	t= 30-425 cm l= 6 to 29 km	GR= 41-308 API DT= 225-355 μsm^{-1} NPHI= 0.17-0.45 v/v RHOB= 2280-2820 gcm^{-3}	Restricted marine shale
	A2	Interbedded claystone and very fine-grained sandstone; non-parallel and wavy lamination. Scarcely bivalve shells oriented parallel to bedding.	t= 10-725 cm l= 8 km to 13 km	GR= 71-65 API DT= 189-268 μsm^{-1} NPHI=? RHOB= 2280-2820 gcm^{-3}	Muddy Shallow bay-fill
	A3	Fine to medium grained sandstone; moderately to well sorted grains. Wavy bedding, cross bedding, rare wave ripples	t= 60-370 cm l < 8 km	GR= 18-46 API DT= 199-314 μsm^{-1} NPHI= 0.07-0.52 v/v RHOB= 1690-2745 gcm^{-3}	Sandy shallow bay-fill
	A4	Coarse to fine-grained sandstones with alternating upward fining to coarsening trend. Moderately sorted grains. Sparse sedimentary structures.	t= 250-500 cm l= 1.8 km to 4.2 km	GR= 7-35 API DT= 175-230 μsm^{-1} NPHI= 0.038-0.146 v/v RHOB= 2280-2820 gcm^{-3}	Marine channel-fill sandstones
B	B1	Upward-coarsening siltstone to fine-grained moderate sorted sandstones, with shell debris, and quartz granules.	t= 30-480 cm l = <2 km	GR= 18-80 API DT= 168-291 μsm^{-1} NPHI= 0.038-0.191 v/v RHOB= 2322-2723 gcm^{-3}	Distal lower shoreface
	B2	Very fine-fine grained, moderate to well sorted sandstone. Fine grained carbonaceous laminae, typically low angle cross beds.	t= 130-440 cm l = 1.7 km – 8 km	GR= 20-56 API DT= 179-277 μsm^{-1} NPHI= 0.048-0.168 v/v RHOB= 2314-2696 gcm^{-3}	Proximal lower shoreface
	B3	Coarsening upward, cross laminated, fine to medium grained, well sorted sandstone; consist carbonaceous fragments	t= 425-800 cm l = 1.7 km – 8 km	GR= 15-25 API DT= 250-275 μsm^{-1} NPHI= 0.09-0.113 v/v RHOB= 2271-2342 gcm^{-3}	Upper Shoreface
C	C1	Highly bioturbated siltstone to very fine sandstones, which has beds of rounded granules	t= 175-1010 cm l = 7.2 km – 19.6 km	GR= 20-80 API DT= 230-260 μsm^{-1} NPHI= 0.08-0.169 v/v RHOB= 2327-2521 gcm^{-3}	Distal mouth bar
	C2	Very fine to fine grained sandstones; low angle cross-bedding.	t= 290-775 cm l = < 5 km	GR= 12-58 API DT= 167-397 μsm^{-1} NPHI= 0.05-0.595 v/v RHOB= 1612-2705 gcm^{-3}	Proximal mouth bar
D	D1	Fining upward coarse to fine grained sandstone; stacked fining upward beds with rare coarse grained stringers.	t= 740-820 cm l = < 2 km	GR= 8-134 API DT= 235-335 μsm^{-1} NPHI= 0.14-0.460 v/v RHOB= 2284-2570 gcm^{-3}	Tidally influenced fluvial channel fill sandstone
	D2	Fining upward coarse to medium grained sandstone. Carbonaceous laminae and fragments. Sharp, and cohesive contact at bed base	t= 580 cm l = < 2 km	GR= 9-34 API DT= 241-297 μsm^{-1} NPHI= 0.14-0.289 v/v RHOB= 2168-2447 gcm^{-3}	fluvial channel fill sandstone
E	E1	Coal and carbonaceous shale. Basal contact, typically parallel.	t= 30-520 cm l = 6 km to 19.6 km	GR= 8-56 API DT= 313-427 μsm^{-1} NPHI= 0.24-0.529 v/v RHOB= 1930-2225 gcm^{-3}	coal
	E2	Alternating dark grey mud/claystone and siltstone to very fine-grained sandstone. Wavy to non-parallel lamination.	t= 60 cm l = < 2 km	GR= 32-60 API DT= 358-415 μsm^{-1} NPHI= 0.43-0.49 v/v RHOB= 1994-2148 gcm^{-3}	Coastal plain fines

Table 2

		Initial Conditions- GPM Input Parameters												
		Simulation Duration	Sediment Type Proportion (%)				Avg. Water Velocity	Avg. Sediment Velocity	Erodibility	Diffusion Coefficient	Avg. Sea Level	Turbidite Event Interval	Steady Flow Iteration	Sediment Movement
GPM Scenarios (GS)		(Ma- 0a) Years	Sand (Coarse)	Sand (Fine)	Silt	Clay	(m/a)	(m/a)			Interval (m)	(/years)	(/hrs)	Coefficient
	S1	0.02 – 0	25	25	25	25	0.11	0.03	0.35	0.11	30	2500	10	0.001
	S2	0.25 – 0	25	25	25	25	0.15	0.03	0.45	0.15	70	1000	15	0.012
	S3	0.5 – 0	25	25	25	25	0.11	0.02	0.55	0.11	120	1000	20	0.012
	S4	0.7 – 0.05	25	25	25	25	0.08	0.02	0.35	0.08	100	500	25	0.0011
	S5	1.5 – 0	15	35	30	20	0.15	0.04	0.50	0.15	80	5000	20	0.001
	S6	3.0 – 0	50	25	15	10	0.13	0.04	0.50	0.13	70	5000	30	0.0012
	S7	3.5 – 0	50	25	15	10	0.11	0.04	0.50	0.11	70	10000	15	0.001
	S8	4.0 – 0	50	25	15	10	0.13	0.04	0.50	0.13	90	5000	20	0.0015
	S9	4.5 – 0	15	45	25	15	0.1	0.02	0.45	0.1	50	10000	30	0.0012
	S10	5.0 – 0	15	45	25	15	0.12	0.02	0.45	0.12	55	10000	35	0.0013
	S11	5.5 - 0	15	45	25	15	0.12	0.02	0.45	0.12	40	5000	40	0.0013
	S12	6.0 – 0	15	45	25	15	0.1	0.02	0.45	0.1	60	10000	35	0.0011
	S13	6.5 – 0	10	25	55	10	0.13	0.03	0.48	0.13	100	20000	50	0.0010
	S14	7.0 – 0	10	25	55	10	0.16	0.03	0.48	0.16	40	20000	45	0.0011
	S15	7.5 – 0	10	25	55	10	0.13	0.03	0.48	0.13	40	20000	40	0.0012
	S16	8.0 – 0	10	25	55	10	0.15	0.03	0.48	0.15	30	10000	30	0.0010
	S17	8.5 – 0	10	25	45	20	0.14	0.02	0.45	0.14	50	50000	50	0.0010
	S18	9.0 – 0	30	30	18	22	0.13	0.02	0.52	0.13	60	25000	35	0.0012
	S19	9.5 – 0	30	40	12	18	0.12	0.02	0.55	0.12	55	25000	20	0.0013
S20	10.0 - 0	30	42	18	10	0.11	0.01	0.40	0.11	50	5000	15	0.0011	
Sediment Property														
	Sediment Type	Diameter	Density	Initial Porosity	Initial Permeability	Compacted Porosity	Compaction	Compacted Permeability	Erodibility					
	Coarse Grained Sand	1.0 mm	2.70 g/cm ³	0.21 m ³ /m ³	500 mD	0.25 m ³ /m ³	5000 KPa	50 mD	0.6					
	Fine Grained Sand	0.1 mm	2.70 g/cm ³	0.3 m ³ /m ³	100 mD	0.15 m ³ /m ³	2500 KPa	5 mD	0.45					
	Silt	0.01 mm	2.65 g/cm ³	0.38 m ³ /m ³	50 mD	0.12 m ³ /m ³	1200 KPa	2 mD	0.3					
	Clay	0.001 mm	2.65 g/cm ³	0.48 m ³ /m ³	5 mD	0.05 m ³ /m ³	500 KPa	0.1 mD	0.15					

Table 3

Lithofacies Classification		
Facies Code	Lithofacies	Command Used in Petrel's Property Calculator
0	Marine Shale	If(Sand_fine>=0.19 And Sand_fine<=0.21 Or Silt>=0.19 And Silt<=0.2 Or Clay>=0.2 And Clay<=0.21 Or Depth_of_deposition>=-82 And Depth_of_deposition<=-78)
1	Muddy Shallow Bay Fill	If(Sand_fine>=0.36 And Sand_fine<=0.38 Or Silt>=0.18 And Silt<=0.2 Or Clay>0.18 And Clay<=0.19 Or Depth_of_deposition>=-30 And Depth_of_deposition<=-20)
2	Sandy Shallow Bay Fill	If(Sand_coarse>=0.65 And Sand_coarse<=0.73 Or Sand_fine>=0.18 And Sand_fine<=0.22 Or Silt>=0.18 And Silt<=0.2 Or Clay>=0.17 And Clay<=0.18 Or Depth_of_deposition>=-3 And Depth_of_deposition<=0)
3	Channel Fill Sandstone	If(Sand_coarse>=0.5 And Sand_coarse<=0.68 Or Sand_fine>=0.23 And Sand_fine<=0.25 Or Silt>=0.17 And Silt<=0.18 Or Depth_of_deposition>=0 And Depth_of_deposition<=2)
4	Lower Shoreface Units	If(Sand_coarse>=0.19 And Sand_coarse<=0.31 Or Sand_fine>=0.19 And Sand_fine<=0.24 Or Silt>=0.4 And Silt<=0.48 Or Clay>=0.19 And Clay<=0.31 Or Depth_of_deposition>=-83 And Depth_of_deposition<=50)
5	Middle Shoreface Units	If(Sand_coarse>=0.32 And Sand_coarse<=0.53 Or Sand_fine>=0.25 And Sand_fine<=0.32 Or Silt>=0.26 And Silt<=0.32 Or Clay>=0.19 And Clay<=0.21 Or Depth_of_deposition>=-38 And Depth_of_deposition<=-12)
6	Upper Shoreface Units	If(Sand_coarse>=0.53 And Sand_coarse<=0.72 Or Sand_fine>=0.28 And Sand_fine<=0.33 Or Silt>=0.16 And Silt<=0.21 Or Depth_of_deposition>=-10 And Depth_of_deposition<=6)
7	Distal Mouth Bar Units	If(Sand_fine>=0.23 And Sand_fine<=0.27 Or Silt>=0.38 And Silt<=0.43 Or Clay>=0.19 And Clay<=0.21 Or Depth_of_deposition>=-95 And Depth_of_deposition<=-80)
8	Proximal Mouth Bar Units	If(Sand_coarse>=0.53 And Sand_coarse<=0.71 Or Sand_fine>=0.27 And Sand_fine<=0.32 Or Silt>=0.16 And Silt<=0.21 Or Clay>=0.06 And Clay<=0.07 Or Depth_of_deposition>=-30 And Depth_of_deposition<=-27)
9	Tide Influenced Sandstones	If(Sand_coarse>=0.53 And Sand_coarse<=0.71 Or Sand_fine>=0.26 And Sand_fine<=0.31 Or Silt>=0.35 And Silt<=0.41 Or Depth_of_deposition>=-5 And Depth_of_deposition<=1)
10	Fluvial Channel Sandstones	If(Sand_coarse>=0.54 And Sand_coarse<=0.56 Or Sand_fine>=0.27 And Sand_fine<=0.29 Or Silt>=0.19 And Silt<=0.21 Or Depth_of_deposition>=-2 And Depth_of_deposition<=2)
11	Coal	Estimated as background attribute
12	Coastal plain fines	If(Silt>=0.31 And Silt<=0.43 Or Clay>=0.31 And Clay<=0.35 Or Depositional_depth>=-100 And Depositional_depth<=-40)
13	Marine Mudstone	If(Sand_fine>=0.36 And Sand_fine<=0.38 Or Silt>=0.4 And Silt<=0.52 Or Clay>=0.45 And Clay<=0.78 Or Depth_of_deposition>=-105 And Depth_of_deposition<=-90)

Table 4

Code	Lithofacies	Average NPHI	Density Porosity	Estimated Porosity	KLOGH (mD)
0	Marine Shale	0.17 - 0.45	0.1	0.08 - 0.11	10.02 - 16.1
1	Muddy Shallow Bay Fill	0.17 - 0.42	0.1	0.08 - 0.13	23.85 - 102.3
2	Sandy Shallow Bay Fill	0.07 - 0.52	0.25	0.16 - 0.25	100.0 - 398.7
3	Channel Fill Sandstone	0.04 - 0.15	0.30	0.18 - 0.22	400.01 - 889.7
4	Distal Lower Shoreface	0.04 - 0.19	0.29	0.1 - 0.23	120.5 - 170.3
5	Proximal Shoreface	0.05 - 0.17	0.31	0.17 - 0.24	80.2 - 412.5
6	Upper Shoreface Units	0.09 - 0.11	0.28	0.21 - 0.26	650.2 - 1023.7
7	Distal Mouth Bar Units	0.08 - 0.17	0.27	0.09 - 0.17	170.5 - 223.1
8	Proximal Mouth Bar	0.05 - 0.59	0.12	0.19 - 0.21	130.5 - 314.3
9	Tide Influenced SS	0.14 - 0.46	0.26	0.15 - 0.20	220.0 - 512.6
10	Fluvial Sandstones	0.14 - 0.29	0.21	0.19 - 0.21	180.5 - 691.8
11	Coal	0.24 - 0.53	0.05	0.001	0.001
12	Coastal Plain Fines	0.43 - 0.49	0.06	0.04 - 0.12	5.2 - 34.6
13	Marine Mudstone	0.16 - 0.42	0.1	0.08 - 0.10	6.0 - 15.2

Porosity and Permeability Prediction through Forward Stratigraphic Simulations Using GPM™ and Petrel™: Application in Shallow Marine Depositional Settings.

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Abstract

The forward stratigraphic simulation approach ~~is was~~ used ~~in this work~~ to ~~predict model~~ porosity and permeability attributes in the Volve field, Norway. This was achieved by ~~using applying~~ spatial data from the forward stratigraphic ~~model simulation~~ to control ~~the property~~ distribution ~~of porosity and permeability~~ in the ~~3-D grid reservoir model~~. Building a ~~subsurface property reservoir~~ model that fits data at different locations ~~in a hydrocarbon reservoir~~ is a task associated with high levels of uncertainty. ~~An appropriate means to~~ To minimise property representation uncertainties ~~is to use~~ in a reservoir model, geologically realistic sediment distribution ~~and or stratigraphic~~ patterns must be developed to predict lithofacies units and ~~related associated~~ petrophysical properties. The workflow ~~used are is~~ in three parts; first, the geological process modeling (GPM™) software developed by Schlumberger was used to simulate scenarios of sediment transportation and deposition in the model area. Secondly, an estimation of lithofacies proportions in the stratigraphic model was done using the property calculator tool in the Petrel™ software. Finally, porosity and permeability values ~~are were~~ assigned to corresponding lithofacies-associations in the forward model to produce a forward stratigraphic-based petrophysical model. Results show a lithofacies distribution that is controlled by sediment diffusion rate, sea level variation, flow rate, wave processes, and tectonic events. This observation is consistent with real-world events, ~~were~~ variation in sea level ~~changes~~, volume of sediment input, and accommodation ~~space~~ control the ~~kind~~ build-up of stratigraphic sequence ~~formed~~. Validation wells ~~prefixes~~, VP1 and VP2 located in the original Volve field petrophysical model and the forward stratigraphic-based models show a good match in porosity and permeability attributes at 5 m vertical sample intervals. ~~By reducing the level of property uncertainty between wells through~~ The resultant forward stratigraphic ~~modeling, an improved based~~ porosity and

permeability models suggest that forward stratigraphic simulation outputs can be achieved for an efficient field development strategy integrated into classical modeling workflows to improve subsurface property representation, and well planning strategies.

1 Introduction

2 The distribution of reservoir properties such as porosity and permeability is a direct function of a complex
3 combination of sedimentary, geochemical, and mechanical processes (Skalinski & Kenter, 2014). The
4 impact of reservoir petrophysics on ~~hydrocarbon field development and depletion~~well planning and
5 extraction strategies makes it imperative to use reservoir modeling techniques that present realistic
6 property variations in 3-D models (~~e.g.~~Deutsch and Journel, 1999; Caers and Zhang, 2004; Hu &
7 Chugunova, 2008). Typically, reservoir modeling ~~tasks~~ require continued property modification until an
8 a appropriate match to known subsurface data is obtained. However, acquisition of subsurface datasets is
9 costly, thus restricts data collection and subsurface modeling ~~condition~~conditions. Several studies, e.g.
10 Hodgetts et al. (2004) and Orellana et al. (2014) have demonstrated that stratigraphic patterns, and
11 therefore petrophysical attributes can be ~~fairly well understood~~extrapolated from seismic, outcrop and
12 well logs. However, this notion is limited by the absence of ~~an~~ accurate and reliable 3-D depositional
13 ~~model~~models to guide ~~the distribution of~~ property variability modeling in reservoir units (Burgess et al.
14 2008). Reservoir modeling techniques with the capacity to integrate forward stratigraphic simulation
15 outputs with stochastic modeling techniques for subsurface property modeling will improve reservoir
16 heterogeneity characterization, because they more accurately produce geological realism than the other
17 modeling methods (Singh et al. 2013). The use of geostatistical-based methods to represent the spatial
18 variability of reservoir properties have been widely accepted in many exploration and production projects
19 (e.g. Kelkar and Godofredo, 2002). In geostatistical-~~base~~ modeling methods, an alternate numerical 3-D
20 model (i.e. realizations) is derived to demonstrate different scenarios of property distribution that can be
21 conditioned to well data (Ringrose & Bentley, 2015). ~~Typically, subsurface~~Reservoir modeling
22 practitioners are normally faced with the challenge of getting a lot of subsurface data to deduce reliable
23 variogram models as a result of cost, therefore introducing a significant level of uncertainty in a reservoir
24 model (Orellana et al. 2014). The advantages of applying geostatistical approaches in populating
25 propoerties in reservoir models is well established (e.g. Deutsch and Journel, 1999; Dubrule, 1998), but

26 the geostatistical-based method tends to confine reservoir property models to known data and rarely
27 realize geological realism to capture sedimentary that have led to reservoir formation (Hassanpour et al.
28 2013). In effect, the geostatistical modeling technique is unable to reproduce a long-range continuity
29 of continuous reservoir properties that are essential for generating realistic reservoir connectivity models
30 (Strebelle & Levy, 2008). Based on lessons from a previous work (e.g. Otoo and Hodgetts, 2019), the
31 forward stratigraphic simulation approach is again was applied in this contribution to predict lithofacies
32 units, porosity, and petrophysical permeability properties in a 3-D model. An important aspect of this work
33 is the use of variogram parameters from forward stratigraphic-based synthetic wells to populate
34 petrophysical properties, especially within inter-well regions of in the reservoir under study model grid.
35 Forward stratigraphic modeling involves the uses morphodynamic rules to derive sedimentary
36 depositional pattern trends to reflect stratigraphic observations patterns in real known data. The approach
37 is driven by the principle that multiple sedimentary process-based simulations in a 3-D framework will
38 most likely improve our understanding on spatial variation of facies, as well as and petrophysical
39 properties property distribution in a geological system model.

40 The sedimentary system, Hugin formation makes up the main reservoir interval in the Volve field.
41 According to studies under study is located within the Hugin formation. Studies by Varadi et al. (1998);
42 Kieft et al. (2011), suggest that the Hugin formation is made up consist of a complex depositional
43 architecture of waves, tidestidal and riverine riverine fluvial processes; suggesting. This indicates that a
44 single depositional model will not be adequate to produce a realistic lithofacies distributions model of the
45 area. Furthermore, the complicated Syn-depositional rift-related faulting system, significantly influence
46 the stratigraphic architecture (Milner and Olsen, 1998). The focus of this study work is to produce a
47 depositional sequence in the shallow marine environment by using a forward stratigraphic modeling
48 approach in the GPMTM (Schlumberger, 2017), and use variogram parameters from the forward model to
49 control porosity and permeability property representation in a 3-D the Volve field model grid.

73 Study Area

74 The Volve field (Figure 1), located in Block 15/9 south of the Norwegian North Sea is Jurassic in age
75 (i.e. late Bajocian to Oxfordian) with the Hugin Formation as the main reservoir unit from which
76 hydrocarbons are produced (Vollset and Dore, 1984). The Hugin formation is made up of shallow marine
77 to marginal marine sandstone deposits, coals, and a significant influence of wave events that tend to
78 control lithofacies distribution in the formation (Varadi et al. 1998; and Kieft et al. 2011). Several studies,
79 e.g. Sneider et al. (1995), and Husmo et al. (2003) associate sediment deposition in the Hugin system to
80 a rift-related subsidence and successive flooding during a large transgression of the Viking Graben within
81 the Middle to Late Jurassic period. Previously it was interpreted to comprise of marine shoreface, lagoonal
82 and associated coastal plain, back-stepping delta-plain and delta front deposits (e.g. Cockings et al. 1992;
83 Milner and Olsen, 1998), but recent studies, e.g. Folkestad and Satur, (2006) suggest the influence of a
84 strong tidal event, which introduces another dimension in property modeling of the reservoir. The
85 thickness of the Hugin formation is estimated to range between 5 m and 200 m but can be thicker off-
86 structure and non-existent on structurally high segments as a result of post-depositional erosion (Folkestad
87 and Satur, 2006).

88 Based on studies by Kieft et al. (2011), a summarised sedimentological delineation within the Hugin
89 formation is presented in **Table 1**. Lithofacies-association codes A, B, C, D, and E used in the
90 classification represents bay fill units, shoreface sandstone facies, mouth bar units, fluvio-tidal channel
91 fill sediments, and coastal plain facies units respectively. In addition a lithofacies association prefixed
92 code F was interpreted to consist of open marine shale units, mudstone with occasional siltstone beds,
93 parallel laminated soft sediment deformation that locally develop at bed tops. The lateral extent of the
94 code F lithofacies package in the Hugin formation is estimated to be 1.7 km to 37.6 km, but the total
95 thickness have of code F lithofacies is not been completely penetrated known (Folkestad & Satur, 2006).

96 **Data and Software**

97 This work is based on description, and interpretation of petrophysical datasets in the Volve field by
98 Statoil, now Equinor. Datasets include 3-D seismic data sections, and a suite of 24 wells that consist of
99 formation pressure data, core data, and sedimentological logs. Previous works such as Folkestad & Satur,
100 (2006) and Kieft et al. (2011) show varying grain size, sorting, sedimentary structures, bounding contacts
101 of sediment matrix that play a significant part of the reservoir petrophysics. Grain size, sediment matrix
102 and the degree of sorting will typically drive the volume of void created, and therefore the porosity and
103 permeability attributes. Wireline-log attributes such as gamma ray (GR), sonic (DT), density (RHOB),
104 and neutron-porosity (NPHI) were used to distinguish lithofacies units, stratigraphic horizons and zones
105 that are required to build the 3-D property model. Porosity, and permeability models, of the Volve field,
106 were generated in Schlumberger's PetrelTM software. Importantly, this work also seeks to produce
107 geologically realistic depositional architecture that is comparable to a real-world stratigraphic framework
108 in a shallow marine environment. Deriving a representative 3-D stratigraphic model of the reservoir
109 allows us to deduce geometrical and variogram parameters as input datasets in actual subsurface property
110 modeling.

111 ~~The Schlumberger's~~ geological process modeling (GPMTM) software ~~developed by Schlumberger~~ was
112 used to undertake twenty forward stratigraphic simulation in an attempt to replicate ~~the~~ depositional
113 processes that resulted in the build-up of the reservoir interval under study. Simulations were constrained
114 to twenty scenarios because the desired stratigraphic sequence and associated sediment patterns were
115 achieved at the fourth simulation. The main criteria for evaluating the realistic nature of a stratigraphic
116 model was to compare it to the depositional sequence observed in the seismic section in Figure 3b. Several
117 process modeling software packages exist and have been applied in similar studies; e.g. Delft3D-FlowTM;
118 Rijin & Walstra, (2003); DIONISOSTM; Burges et al. (2008). The geological process modeling (GPMTM)
119 software was preferred because of the availability of software license, and also the ease in integrating of
120 its outputs into the property modeling workflow in PetrelTM.

147 Methodology

148 The workflow (Figure 2a) combines the stratigraphic simulation capacity of the GPM™ software in
149 different depositional settings, and the property modeling tools in Petrel™ to predict the distribution of
150 porosity and permeability properties away from well data. Three broad steps have been used here to
151 achieve this goal; (i) forward stratigraphic simulation (FSS) in GPM™ software (2019.1 version), (ii)
152 lithofacies classification using the calculator tool in Petrel™, and (iii) lithofacies, porosity, and
153 permeability modeling in Petrel™ (2019.1 version).

154 Process Modeling in GPM™

155 The GPM™ software consist of different geological processes that are designed to replicate sediment
156 deposition in clastic and carbonate environments. ~~Example, the steady flow process is efficient for~~
157 ~~simulating sediment deposition in fluvial bodies, whilst the unsteady flow process control sediment~~
158 ~~transportation from the basin slope into deep water basin setting, largely in the form of basinal floor fan~~
159 ~~units. Previous~~For example, previous studies, e.g. Kieft et al, (2011) identified the influence of riverine
160 (fluvial), and wave processes in the genetic structure of sediments in the Hugin formation. These
161 geological processes could be very rapid depending on accommodation ~~space~~ generated as a result of sea
162 level variation, and or sediment composition and flow intensity. Sediment deposition, and its response to
163 post-depositional sedimentary and tectonic processes are significant in the ultimate distribution of
164 subsurface lithofacies units; ~~hence the variation of input parameters to increase our chance attaining. To~~
165 ~~attain stratigraphic~~ outputs that fall within ~~acceptable limits of what may exist~~ the depositional architecture
166 captured in the ~~natural order seismic section (Figure 3b), the input parameters were varied as illustrated~~
167 by different scenarios in Table 2. The simulation generated geologically realistic stratigraphic
168 ~~frameworks~~ trends, but also revealed some limitations, such as instability in the simulator when more than
169 three geological processes and sub-operations run at a time. In view of this, the diffusion and tectonic
170 processes ~~are constant features whiles~~ were combined with other processes ~~likesuch as~~ steady flow,
171 unsteady flow, and sediment accumulation, ~~compaction were varied to replicate the Volve field~~
172 stratigraphic depositional scenarios.

Steady Flow Process

The steady flow process in GPM simulate flows that changes slowly over a period, or sediment transport scenarios where flow velocity and channel depth do not vary abruptly; e.g. rivers at normal stage, deltas, and sea currents. The steady flow process can be specified to a desired setting in the “run sedimentary simulation” dialog box in the Petrel™ software (version 2017.1 and above). Considering the influence of fluvial activities in the build-up of the Hugin formation, it was important to capture its impact on the resultant simulated output. To attain stability in the simulator, it is advisable to undertake preliminary runs to ascertain the appropriateness of input parameters that will be used in the simulation. For steady flow process, a boundary condition must be specified at the edges of the model. For example in an open flow system, negative integers (i.e. values below zero) must be assigned to the edges of the hypothetical paleo-surface to allow water to enter and leave the simulation area.

Unsteady Flow Process

The unsteady flow process can model flows that are periodic, and run for a limited time; for example, in turbidites where velocity of flow and depth changes abruptly over time. The unsteady flow process algorithm is set up to apply a number of fluid elements, that are affected by gravity, and by friction against the hypothetical topographic surface. A contribution on the application of the unsteady in stratigraphic simulation, and how its settings can be configured to attain geological realistic outcomes is discussed in Otoo and Hodgetts, (2019).

Diffusion Process

The diffusion process can effectively replicate sediment erosion from areas of higher slope (i.e. source location) and their deposition to lower elevation of the model area. Sediment dispersion in the diffusion process is carried out through erosion and transportation processes that are driven by gravity in the simulator. The diffusion process is based on the assumption that sediments are transported downslope at a proportional rate to the topographic gradient; therefore making fine grained sediments easily transportable than coarse grained sediments. Diffusion is controlled by two parameters; (i) diffusion

223 coefficient, which controls the strength of the diffusion, and (ii) diffusion curve that serves as a unitless
224 multiplier in the algorithm. The governing equation for the diffusion process is:

225 $\frac{\partial z}{\partial t} = k \nabla^2 z$, where z is topographic elevation, k the diffusion coefficient, t for time, and $\nabla^2 z$ is the
226 laplacian.

227 **Sediment Accumulation**

228 In the GPMTM software, sediment source can be set to a point location or considered to emanate from a
229 whole area. Sediment accumulation deals with sediment deposition via an areal source. For example,
230 where a lithology is interpreted to be uniformly distributed, the sediment accumulation process can be
231 used to replicate such depositional scenario. The areal input rates for each sediment type (e.g. coarse
232 grained, fine grained sediments) used in the accumulation process must be specified in the settings.
233 Specifying the areal rates for each sediment is important because the software is configured to use the
234 value of the surface at each cell in the model grid and multiplies it by a value (i.e. value from a unitless
235 curve) at each time step in the simulation to estimate the thickness of sediments accumulated or eroded
236 from a cell in the model.

237 **Parameters for Forward Stratigraphic Simulation**

238 A realistic reproduction of stratigraphic patterns in the study area will require input parameters (~~also~~
239 ~~known as or~~ initial conditions). ~~These include: a such as:~~ hypothetical paleo-topography, sea level curves,
240 sediment source location and distribution curve, tectonic ~~event~~ event maps (i.e. subsidence and uplift),
241 and sediment mix velocity. The application of these input parameters in ~~the~~ GPMTM ~~simulator~~, and their
242 influence on the resultant stratigraphic framework are ~~explained~~ discussed below.

243 **Hypothetical Paleo-Surface:** The hypothetical paleo-topographic surface, on which the simulation
244 ~~commence~~ evolves was inferred from the seismic section. ~~Here, we assume~~ This is done with the assumption that
245 the present day stratigraphic surface, ~~also referred to as the~~ (i.e. paleo shoreline in Figure 3a) occurred as a result
246 of basin filling through different geological periods. Since the hypothetical topography generated from the seismic
247 section have undergone various phases of subsidence and uplifts ~~over time~~, the paleo topographic surface used in

274 this work does not present an accurate description of the basin at the period of sediment deposition. To
275 ~~mitigate~~obtain an appropriate paleo-topographic for this ~~uncertainty, 5task, , five~~ paleo topographic
276 surfaces (TPr) were generated ~~stochastically~~ by adding or subtracting elevations from the inferred paleo
277 topographic surface ~~or base topography~~ (see Figure 4g) using the equation: $TPr = Sbs + EM$, where, Sbs is
278 the base surface scenario (in this instance, scenario 6), and EM an elevation below and above the base
279 surface. ~~In this work,~~Paleo-topographic surface in scenario 3 (figure ~~3d4d~~) was ~~used as the paleo topographic~~
280 ~~surfaces~~selected, because it ~~produced~~controlled the development of stratigraphic sequences that fit the
281 conceptual knowledge of depositional framework as observed in the seismic section (Figure 5d).

282 **Sediment Source Location:** Based on regional well correlations in previous studies (e.g. Kieft et al.
283 2011), and ~~seismic interpretation of~~ the basin structure ~~interpreted from seismic data~~, the sediment entry
284 point ~~for this task~~ was placed in the north-eastern section of the hypothetical paleo-topography. Since the
285 exact sediment entry point is ~~uncertain~~not known, multiple entry points were placed at 4 ~~m~~km radius
286 around the primary location in (Figure 3c), in order to capture possible sediment source locations. The
287 source position is characterised by positive integers (i.e. values greater than zero) to enable fluid flow to
288 other parts of the simulation surface.

289 **Sea Level:** ~~Primarily, the sea~~Sea level variation ~~relative to elevation~~ was inferred from published studies
290 and facies description in shallow marine depositional environments (e.g. Winterer and Bosellini, 1981).
291 ~~Considering the limitations~~To attain stability in the ~~software~~simulator, we assumed a sea level ~~of that~~
292 ~~range between 15 m to 45 m; averaging~~ 30 m for short simulation runs, e.g. ~~5000 to~~ 20000 years ~~to attain~~
293 ~~stability in the simulator and.~~ The sea level was varied ~~it accordingly~~ with increasing duration of the
294 simulation. ~~(illustrated in Table 2).~~ The peak sea-level in the simulation represents the maximum flooding
295 surface, ~~(Figure 5d),~~ and therefore ~~an~~the inferred sequence boundary in the geological process model.

296 **Diffusion and Tectonic Event Rates:** The sediment mix proportion ~~and,~~ diffusion rate ~~for the simulation~~
297 ~~and tectonic event functions~~ were ~~stochastically~~ inferred from previous studies (e.g. ~~Burges et al., 2008~~),
298 ~~primarily to attain a prograding and or aggrading clinoforms features that are noticeable in real world~~
299 ~~geological outcrops. The subsidence and uplift rates were kept constant in most part of the model. The~~

300 ~~functions are inferred from published works; e.g.~~ Walter, 1978; Winterer and Bosellini, 1981, and Burges
301 et al., 2008). ~~The diffusion and tectonic event rates are~~ increased or reduced to produce a stratigraphic
302 model that fit our knowledge of the basin evolution. ~~The simulation parameters applied (Table 2) were~~
303 ~~generated randomly using the initial run (Figure 6a) as a guide. The guiding principle for parameter~~
304 ~~selection is their capacity to produce stratigraphic outputs that depict different depositional scenarios in~~
305 ~~the shallow marine setting. A key criteria for selecting parameters is their capacity to produce stratigraphic~~
306 ~~outputs that depict depositional scenarios in the study area. For example, in scenario 1 (Figure 6a), the~~
307 ~~early stages of clinoform development show resemblance to interpreted trends in the seismic section~~
308 ~~(Figure 3b). As a result, input figures that were higher and lesser than those used in generating scenario~~
309 ~~1 were generated to serve as the simulation parameters for the twenty scenarios. In scenario 1, a diffusion~~
310 ~~coefficient of 8 m²/a was used to produce a realistic clinoform build-up, so the figure was varied with +/-~~
311 ~~5 to obtain figures that could improve the model derived in scenario 1. The initial topography (TP_r) was~~
312 ~~kept constant throughout a simulation, but wave events, steady/unsteady flow, diffusion and tectonic~~
313 ~~events use curve functions to provide variations within the simulation.~~ A sudden change in subsidence
314 rate tends to constrain coarse to medium sediments at proximal distance to source location than in
315 scenarios where the rate of subsidence was made gradual.

316 The influence of input parameters in the simulation is evident whenever there is a slight change of value
317 in sediment diffusion, and tectonic rates or dimension of the hypothetical topographic surfaces. For
318 example, sediment source position has a strong impact on the extent and depth to which sediments are
319 deposited in the basin. Shifting the source point to the mid-section of the topography (i.e. the mid-point
320 of the topography in a basin-ward direction) resulted in the accumulation of distal elements that are
321 identical to turbidite lobe systems. This is consistent with morphodynamic experiments ~~(e.g. by~~ de Leeuw
322 ~~et al., 2016),~~ where abrupt sediment discharge ~~of sediments~~ from the basin slope leads to the build-up of
323 basin floor fan units. ~~Stratigraphic patterns generated using different input parameters provides 3-D~~
324 ~~perspective into subsurface property variations under alternating initial conditions.~~

Property Classification in Stratigraphic Model

In our opinion, the most appropriate model in this work is **Figure 5d**. This is because, when compared to depositional description in studies such as Folkestad and Satur (2006); Kieft et al., (2011), it produced a stratigraphic sequence that mimics the depositional sequence in the shallow marine depositional environment under study area. The stratigraphic model was converted into a 3-D format, 20 m x 20 m x 2 m grid cells in order to be used in the property modeling tool in Petrel™. Lithofacies, porosity, and permeability properties are characterized in the stratigraphic using a rule based approach (**Table 3**). Sediment distribution in each time step of the simulation were stacked into a single zone framework to attain a simplified model. This was done with the assumption that sedimentary processes that lead to the final build-up of genetic related units within zones of the forward stratigraphic architecture model will not vary significantly over the simulation period. Property classification in the model was achieved with the property calculator tool in Petrel™. The classification is was driven by depositional depth, geologic flow velocity, and sediment distribution patterns as indicated in **Figure 7**. Lithofacies representation in the stratigraphic model was based on the sediment grain size pattern, and proximity to sediment source. For example, shoreface lithofacies units were characterized using medium-to-coarse grained sediments to that are, which accumulate at proximal distance to the sediment source, while. In contrast, mudstone units are constrained to the distal parts of the stratigraphic model, where were restricted to fine grained sediments that accumulate at the end distal section of the simulation domain.

Porosity and Using published studies by Kieft et al., (2011), porosity and permeability variations in the stratigraphic model were estimated from published wireline-log attributes (e.g. Kieft et al., 2011), which is such as gamma ray, neutron, sonic, and density logs outlined in Table 1. Based In previous studies on petrophysical report of the Sleipner Øst, and Volve field (Statoil e.g. Equinor, 2006), a deduction was made; Kieft et al., 2011), Shoreface deposits were identified, to the effect that high net to gross zones will be associated with make up the best quality reservoir units; classified as, whilst lagoonal deposits formed the worst reservoir units. Using this as guide, shoreface lithofacies units, whilst low net to gross zones were interpreted to be connected with high proportions of shale or sandstone units and mudstone

400 ~~deposits/shale units in the forward stratigraphic model were characterized as best and worst reservoir~~
401 ~~units respectively.~~ The porosity and ~~184~~ permeability values in Table 4 were derived from equations in
402 Statoil's petrophysical report of the Volve ~~185~~ field (StatoilEquinor, 2016):

403 $\phi_{er} = \phi_D + \alpha \cdot (NPHI - \phi_D) + \beta$; where ϕ_{er} is the estimated porosity range, ϕ_D is density porosity, α and
404 β are regression constants; ranging between -0.02 – 0.01 and 0.28 – 0.4 respectively, $NPHI$ is neutron
405 porosity. In instances where $NPHI$ values for lithofacies units is not available from the published
406 references, an average of 0.25 was used.

407 $KLOGH_{er} = 10^{(2+8 \cdot PHIF - 5 \cdot VSH)}$; where $KLOGH_{er}$ is the estimated permeability range, VSH is the volume
408 of clay/shale in the lithofacies unit, and $PHIF$, the fractured porosity. The VSH range between 0.01 – 0.12
409 for the shoreface units, and 0.78 – 0.88 for lagoonal deposits.

410 Property Modeling in Petrel™

411 The workflow (Figure 2b) used for subsurface property (e.g. lithofacies, and petrophysical) modeling in
412 Petrel™ is extended to represent lithofacies, porosity, and permeability properties in the forward
413 stratigraphic model. These processes include:

- 414 1. Structure modelling: identified faults within the study area are modelled together with
415 interpreted surfaces from seismic and well data to generate the main structural framework,
416 within which the entire property model will be built. The procedures involve modification of
417 fault pillars and connecting fault bodies to one another to attain the kind of fault framework
418 interpreted from seismic and core data.

- 419 (2) Pillar gridding: a “grid skeleton” that is made up of a top, middle and base architectures. Typically,
420 ~~there are~~ pillars ~~which~~ join corresponding corners of every grid cell of the adjacent ~~grid,~~ forming
421 ~~to form~~ the foundation ~~offor~~ each cell within the model; ~~hence its nomenclature as a corner point~~
422 ~~gridding.~~ The prominent orientation of faults (i.e. I-direction) within the model ~~is~~ area generally

trends in a N-S and NE-SW direction, so the “I-direction” was set to the major NNE-SSW direction along which grid cells align to capture the structural description.

(3) Horizons, Zones and Vertical Layering: stratigraphic horizons and subdivisions (zones) delineates the geological formation’s boundaries. As stratigraphic horizons are inserted into the model grid, the surfaces are trimmed iteratively and modified along faults to correspond with displacements across multiple faults. Vertical layering on the other hand defines the thicknesses and orientation between the layers of the model. Layers in this context describes significant changes in particle size or sediment composition in a geological formation. Using a vertical layering scheme makes it possible to honour the fault framework, pillar grid and horizons that have been derived. A constant cell thicknesses are of 1 m across the model was defined to control the vertical scale, in which subsurface properties such as lithofacies, porosity, and permeability attributes are modelled.

(4) Upscaling, which involves averaging the substitution of finer grid cells in order to with coarser grid cells. This is done to assign property values to the cells and in order to evaluate which discrete value suits each a selected data point. It also encompasses One advantage of the generation of coarser grids (i.e. lower resolution grids) in the geological model, in order upscaling procedure is to make simulation the modeling process faster.

Porosity and Permeability Modeling

The ~~Volve field~~ porosity and permeability model that was built by Equinor for their operations in the Volve field was adopted as the base model. The model, which cover an area of 17.9 km² was generated with the reservoir management software (RMS), developed by Irap and Roxar (Emerson™). The ~~original~~ petrophysical model has a grid dimension of 108 m x 100 m x 63 m, and was compressed by 75.27% of cell size- from an approximated cell size 143 m x 133 m x 84 m. To achieve a comparable model resolution ~~to~~ as the ~~original~~ Volve field porosity and permeability model, the forward stratigraphic output, which had an initial resolution of 90 m x 78 m x 45 m was upscaled to a cell size grid of 107 m x 99 m x 63 m. Two options were explored with respect to the use of variogram parameters derived from forward model-based

475 synthetic wells. Option 1 was to assign porosity and permeability values to the synthetic lithofacies wells
476 ~~to that~~ correspond to known facies-associations as indicated in **Table 4**. The synthetic wells with porosity
477 and permeability data are placed in-between ~~actual well~~ (known data) locations to guide porosity and
478 permeability property distribution in the model. For option 2, the best-fit forward stratigraphic model was
479 populated with porosity, and permeability attributes using the major stratigraphic orientation captured in
480 the seismic data (i.e. NE-SW; 240°) to control property distribution trends. Porosity and permeability
481 were populated into the model by using the property modeling process in Petrel™. Porosity and
482 permeability synthetic logs are then extracted from the forward stratigraphic output to build the porosity
483 and permeability models (**Figure 8**). ~~The second option provided a broader framework for evaluating the~~
484 ~~reliability of forward stratigraphic simulation on property distribution in areas of sparse data.~~ Taking into
485 account the possibility that vertical trends in options 1 and 2 will ~~most likely produce a~~ similar trend in
486 a sampled interval, it is our opinion that option 2 will provide a viable 3-D representation of property
487 variations in the major and minor directions of the forward stratigraphic model. Ten synthetic wells, (SW),
488 ranging between 80 m and a 120 m in total depth (TD) were positioned in the forward model to capture
489 the vertical distribution of porosity-permeability at different sections of the stratigraphic model.
490 ~~Typically, sediment distribution, and associated petrophysical attributes are directly related to depth~~
491 ~~within the geological model; thus aiding in the analysis of the most likely proportions of subsurface~~
492 ~~properties that match with observations in known well data.~~

493 The forward-based synthetic wells (**Figure 9 c**) with porosity and permeability logs were upscaled to
494 ~~populate the original structural model using the sequential Gaussian simulation method.~~ populate the
495 original structural model using the sequential Gaussian simulation method. Here, the synthetic wells
496 derived from the stratigraphic model is to provide an additional well data for use in a traditional modeling
497 workflow as was the case in the building of original Volve model. Considering the advantages of
498 variogram-based modeling in relation to data conditioning, the idea was to get more wells into the model
499 grid to control porosity and permeability distribution. Upscaling the synthetic well data in this context is

525 to “transform” the data from 1-D into a 3-D framework to build the property model. Using the same
526 structural model was to attain a comparable framework for evaluating the modeling outputs.

527 The variogram model (**Figure 10**), of dominant lithofacies units in the formation served as a guide in the
528 estimation of variogram parameters from the forward model. A major and minor range of 1400 m and
529 400 m respectively, and an average sill value of 0.75 derived from forward stratigraphic-based synthetic
530 wells were used to populate porosity and permeability properties in the model. Porosity models were
531 derived with a normal distribution, whilst the permeability models were produced using a log-normal
532 distribution and the corresponding porosity property for collocated co-kriging. Out of fifty model
533 realizations, six realizations that showed some similarity to the original petrophysical model are presented
534 (**Figure 11**). This was accomplished through visual and statistical comparison of zones in the original
535 Volve field model, and the stratigraphic-based porosity/permeability models. The statistical approach
536 involved a comparison of summary statistics from the original Volve model, and the porosity/permeability
537 model generated through forward stratigraphic modeling. The visual comparison on the other hand looked
538 at how geological realistic the output is, and if it conforms with our conceptual idea of the Volve field
539 model.

540 **Results**

541 The stratigraphic model in stage 4 (**Figure 5d iv**) shows the final geometry after 700, 000 years of
542 simulation time. Initial simulation produced a progradation sequence with foreset-like features (**Figure**
543 **5d i**). A sequence boundary, which indicates the highest sea level in the model separates the initial
544 simulated output from the next prograding phase (**Figure 5d ii**). Initiation of an aggradation stacking
545 pattern starts, and becomes prominent in stage 3 (**Figure 5d iii**). ~~This is~~ These sequences are consistent
546 with real-world scenario where sediment supply matchup with accommodation ~~space generated~~ created as
547 a result of the relative ~~constant~~ sea level rise within a period: (e.g. Muto and Steel, 2000; Neal and Abreu,
548 2009). The diffusion process in GPMTM was used to define the stratigraphic architecture before
549 introducing additional geological processes such as steady flow, unsteady flow, wave events to capture

576 the range of possible depositional styles that have been discussed in published literature (e.g. Folkestad
577 & Satur, 2006; Kieft et al., 2011).

578 The impact of the stratigraphic simulation on porosity and permeability representation in the model ~~is was~~
579 evaluated by comparing its outcomes to the ~~original Volve field~~ porosity and permeability models ~~of the~~
580 ~~Volve by~~ using two synthetic well ~~prefixed,~~ VP1 and VP2. ~~The synthetic well are, which were~~ sampled
581 ~~vertically~~ at a 5 m intervals ~~vertically to estimate the distribution of porosity and permeability attributes~~
582 ~~along wells.~~ Considering that the ~~original porosity and permeability Volve field petrophysical~~ model
583 (Figure 11a) have undergone phases of history matching to ~~enable enhance~~ well planning, and ~~guide~~
584 production strategies in the Volve field, it is reasonable to assume that porosity and permeability
585 distribution in the ~~Volve field~~ petrophysical model will be geologically realistic and less uncertain. A
586 good match in porosity was observed in validation wells that penetrate the model realizations; R14, R20,
587 R26, R36, R45, and R49 (Table 5a). ~~The vertical distribution (Figure 12) of shows the~~ porosity ~~variation~~
588 ~~(0.18 – 0.24) in some selected model realizations shows.~~ This value (i.e. 0.18 – 0.24) is within the range
589 ~~of porosity estimates in the Volve field (Equinor, 2016). In view of the limitation in making variations~~
590 ~~within a modal distribution range (i.e. 0.18 – 0.24) that is consistent with the original model. The simulation~~
591 ~~run in GPM™, the~~ forward stratigraphic-based model ~~have been (FSM) was~~ derived with an assumption
592 that variogram parameters, stratigraphic inclination within zones will remain constant. ~~However As a~~
593 ~~result,~~ the original petrophysical model ~~takes into account, which involve~~ other measured attributes;
594 ~~which could be within~~ the ~~main driver of stratigraphic zone was not considered in~~ the ~~differences in~~
595 ~~forward stratigraphic modeling-based permeability estimates model, hence the major variations~~ noted in
596 Table 5b. Typically, a petrophysical model like the Sleipner Øst and Volve field model will take into
597 account other sources of data ~~such as.~~ ~~For example, data from a~~ special core analysis (SCAL) ~~and other~~
598 ~~petrophysical evaluations from will improve~~ the reservoir ~~section, so petrophysics assessment.~~
599 ~~Considering that the FSM approach did not involve these additional information from the formation,~~ it is
600 ~~reasonably reliable~~ ~~reasonable~~ to suggest that the forward stratigraphic-based porosity and permeability
601 models have been adequately conditioned to known subsurface data.

628 Discussion

629 ~~The results~~Results show the influence of sediment transport rate, (~~or in this example,~~ diffusion rate),
630 initial basin topography and ~~proximity to~~ sediment source location on stratigraphic simulation in the
631 GPMTM software. ~~Notably,~~Similar to other studies (e.g. Muto & Steel, 2000; Neal & Abreu, 2009), a
632 variations in sea level controls the volume of sediment that could be retained or transported further into
633 the basin; therefore controlling the kind of stratigraphic sequences that are generated. In a related work
634 by Burges et al. (2008), it was established that;~~for example,~~ a sediment-wedge topset width was directly
635 linked to the initial bathymetry, in which the sediment-wedge structure was formed, as well as the
636 correlation between sediment supply and accommodation rate. This is in line with observations in this
637 ~~workstudy,~~ where the initial sediment deposit control the geometry of subsequent phase of depositions.
638 ~~Since the~~The uncertainty of initial conditions ~~ofused in~~ this ~~basin is uncertain,~~ work led to the generation
639 of multiple simulationforward stratigraphic scenarios ~~were carried out~~ to account for the range of
640 bathymetries that may have influenced ~~the build-up of sediments~~sediment transportation to form the
641 present day Hugin formation. The simulation produced well defined ~~clineforms~~sloping depositional
642 surfaces in a stratigraphic architecture (i.e. clinoforms) and sequence boundaries that depict the pattern
643 observed in the seismic data. ~~As indicated~~Indicated in ~~otherprevious~~ studies, (e.g. Allen and Posamentier,
644 1993; Ghandour and Haredy, 2019) sequence stratigraphy is vital in the characterization of lithofacies in
645 ~~shallow marine settings; hence,~~sedimentary systems. Therefore, a reproduction of stratigraphic sequence
646 in 3-D, using the forward stratigraphic ~~simulation outputs~~modeling approach in GPMTM provide a good
647 framework to ~~better understand the variation of lithofacies units in the~~analyse property variations in a
648 reservoir ~~through a 3-D perspective.~~ A porosity-permeability model ~~that match~~matching the original
649 petrophysical model was produced using synthetic porosity and permeability logs from the forward
650 stratigraphic model as input datasets in the sequential Gaussian simulation algorithm. As mentioned
651 previously, this exercise did not take into account variations in the layering scheme that develops in
652 different zones of the stratigraphic model. Under this circumstance, we concede that there is a possibility
653 to overestimate and or underestimate porosity and permeability properties as observed in some sampled

654 intervals of the validation wells. In view of this, it is our suggestion that forward stratigraphic simulation
655 outputs should be applied as additional data to understand sediment distribution patterns, and associated
656 vertical and horizontal petrophysical trends in the depositional environment than using its outputs as an
657 absolute conditioning data in subsurface property modeling.

658 ~~The assumptions~~Assumptions made ~~in~~with respect to the type of geological processes, and input
659 parameters to use in the simulation ~~significantly~~certainly differ from what ~~may have~~ existed during the
660 period of deposition. ~~Applying~~So, applying stratigraphic models that fit a basin scale description to a
661 smaller scale reservoir context presents another degree of uncertainty in the approach used here. For
662 example, in their study, Burges et al., (2008) shows that the diffusion geological process fits the
663 description of large scale sediment transportation; suggesting that an extrapolation of its outputs into a
664 well-scale framework could produce results that deviate from the real world ~~architecture~~distribution. In
665 reality, sediment deposition into a geological basin is also controlled by mechanical and geochemical
666 processes, ~~which that~~ tend to modify a formations petrophysical attributes (Warrlich et al. 2010), ~~hence,~~
667 ~~the application of~~). Therefore, using different geological processes and initial conditions to ~~produce~~
668 ~~different~~generate depositional scenarios, ~~from which is~~ a ~~best fits stratigraphic framework of the reservoir~~
669 ~~can be selected. Many forward stratigraphic reasonable approach. However, based subsurface modeling~~
670 ~~studies (e.g. on the approach limitation, which are also discussed in similar works (e.g. . Bertoncello et~~
671 ~~al. 2013; Aas et al. 2014; and Huang et al. 2015), have identified and discussed some limitations with the~~
672 ~~technique. Considering that similar challenges were faced in this work,2015)~~ caution must be taken in
673 using ~~theits~~ outputs ~~from forward stratigraphic simulations~~ in real reservoir modeling; as ~~thisit~~ could lead
674 to an increase ~~uncertainty in theproperty~~ representation ~~of lithofacies and petrophysical properties. bias.~~

675 The correlation between reservoir lithofacies and petrophysics have been examined in previous studies,
676 e.g. Falivene et al. (2006) Hu and Chugunova, (2008), but the difference in predicted and actual reservoir
677 character is less understood. This in large part is due to the absence of a realistic 3-D stratigraphic
678 framework to guide reservoir property representation in geocellular models. It is our opinion that forward
679 stratigraphic modeling methods provide reservoir modeling practitioners a better platform to generate

680 appropriate 3-D lithofacies models to improve petrophysical property prediction in a reservoir, but its
681 outputs should be used cautiously and together with verifiable subsurface patterns from seismic and well
682 ~~data~~datasets.

683 **Conclusion**

684 In this paper, spatial data from a forward stratigraphic simulation is combined with subsurface data from
685 the Volve field, Norway to constrain porosity and permeability distribution in inter-well regions of the
686 model area. As caution, the forward stratigraphic simulation scenarios presented in this contribution do
687 not ultimately prove that spatial and geometrical data derived from stratigraphic modeling can be used as
688 absolute input parameters for a real-world reservoir modeling task. Uncertainties in the choice of initial
689 condition and processes for the stratigraphic simulation led the variation of input parameters in order to
690 attain a depositional architecture that is geologically realistic and comparable to the stratigraphic
691 correlation suggested in some published studies of the study area. ~~Significantly, the~~The good match
692 obtained from validation wells in the original and stratigraphic-based petrophysical model, leads us to the
693 suggestion that an integration of variogram parameters from ~~real~~-well data and forward stratigraphic
694 simulation outputs will improve property prediction in inter-well zones. This suggestion is supported by
695 the idea that more conditioning data (well data) will increase the chance of producing realistic property
696 distribution in the model area. In addition, this work also made some key findings:

- 697 1. For a specific application of forward stratigraphic modeling in GPMTM and a range of model
698 parameters, the process of sediment deposition is influenced by diffusion rate, and proximity to
699 sediment source. This is consistent with several published works on sequence stacking and or
700 system tracts in shallow marine settings, but further work with different stratigraphic modeling
701 simulators could be useful in mitigating some of the challenges faced in this work.
- 702 2. A geologically viable 3-D lithofacies distribution in the shallow marine Hugin formation was
703 achieved, which is evident in scenarios where sediment distribution vertically matches with
704 lithofacies variation in a sampled interval in an actual well log.

718 Geologically feasible stratigraphic patterns generated in the forward stratigraphic model provide
719 additional confidence in the representation of lithofacies, and therefore porosity and permeability
720 property variations in the depositional setting under study. ~~By reducing the level of property
721 uncertainty between wells, a reliable reservoir model can be generated to guide field planning and
722 development in the hydrocarbon exploration and production industry.~~

~~723 Future studies will focus on using an artificial neural network approach to classify lithofacies associations
724 in the forward stratigraphic model in an attempt to reduce uncertainties that arise from cognitive or
725 sampling biases in the calculator (or rule based) approach for estimating lithofacies proportion in a
726 forward stratigraphic model. In addition, efforts will be made in a future contribution to compare the
727 stratigraphic property distribution with ones that are generated more classical methods such as sequential
728 indicator simulation (SIS), and object based modeling. The resultant forward stratigraphic-based porosity
729 and permeability model suggests that forward stratigraphic simulation outputs can be integrated into
730 classical modeling workflows to improve subsurface property modeling , and well planning strategies.~~

731 **Data and Code Availability**

732 The datasets used in this work was obtained from Equinor on their Volve field operations, Norway. This
733 include: 24 suits of well logs, and 3-D reservoir models in Eclipse and RMS formats. The data, models
734 (eclipse and RMS formats), and the rule-based calculation script to generate lithofacies and porosity/permeability
735 proportions are archived on Zenodo as Otoo & Hodgetts, (2020).

736 **GPM™ Software**

737 The version (2019.1) of GPM™ software was used in completing this work after an initial 2018.1 version. Available
738 on: <https://www.software.slb.com/products/gpm>. The software license and code used in the GPM™ cannot be
739 provided, because Schlumberger does not allow the code for its software to be shared in publications.

740 **Model Availability in Petrel™**

741 Petrel™ software (2017.1) was initially used for the task, but completed with Petrel™ software (2019.1);
742 available on: <https://www.software.slb.com/products/petrel>. The software run on a windows PC with the
743 following specifications: Processor; Intel Xeon CPU E5-1620 v3 @3.5GHz 4 cores-8 threads, Memory;
744 64 GB RAM. The computer should be high end, because a lot of processing time is required to execute a
745 task. The forward stratigraphic models are achieved in Zenodo as Otoo & Hodgetts, (2020).

746 **Author Contribution**

747 Daniel Otoo designed the model workflow, conducted the simulation using the GPM™ software, and
748 evaluated the results. David Hodgetts converted the Volve field data into Petrel compactible format for
749 easy integration with outputs from the stratigraphic simulation.

750 **Acknowledgement**

751 Thanks to Equinor for making available the Volve field dataset. Also, thanks to Schlumberger for
752 providing us with the GPM™ software license. A special thanks to Schlumberger for providing the

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754 Ghana National Petroleum Corporation (GNPC) for sponsoring this research.

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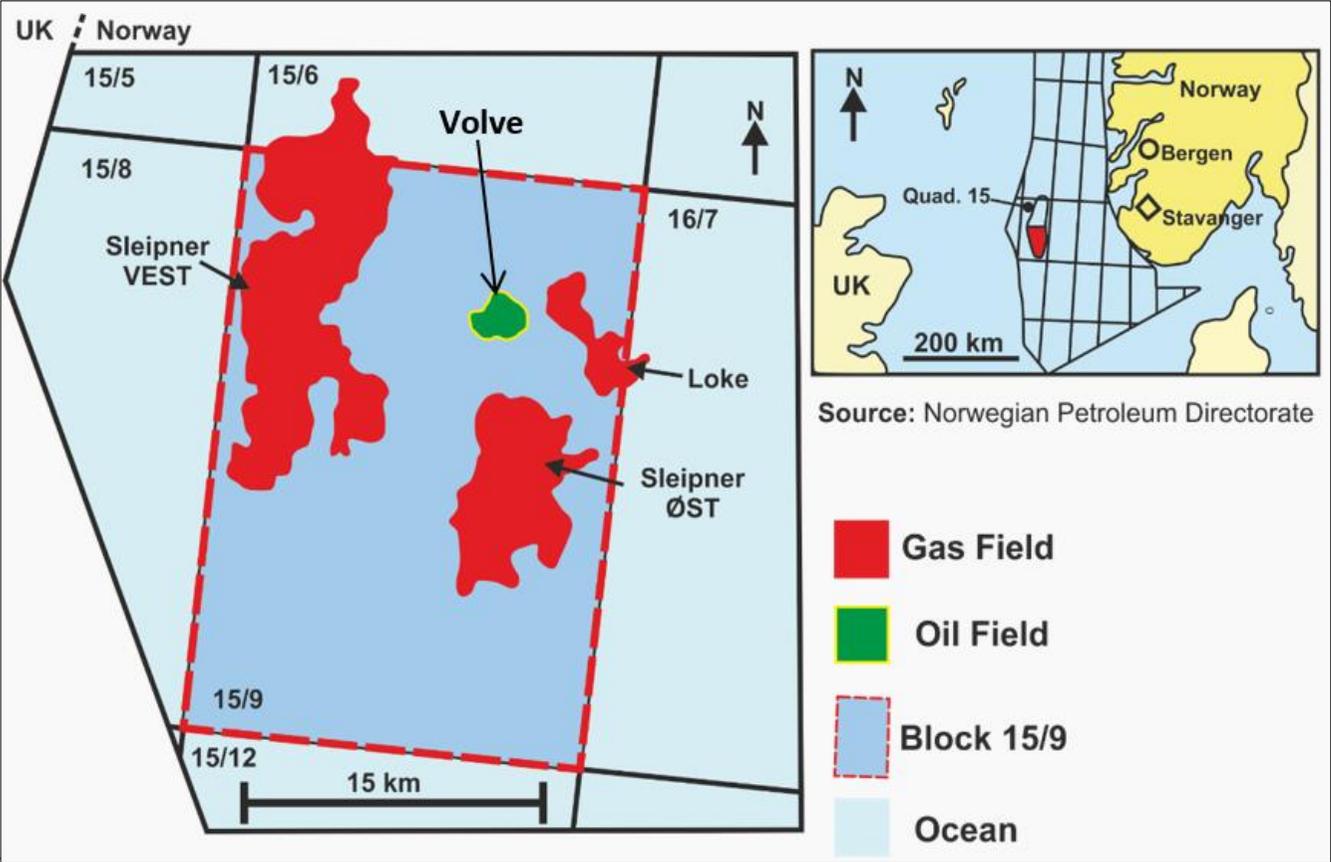
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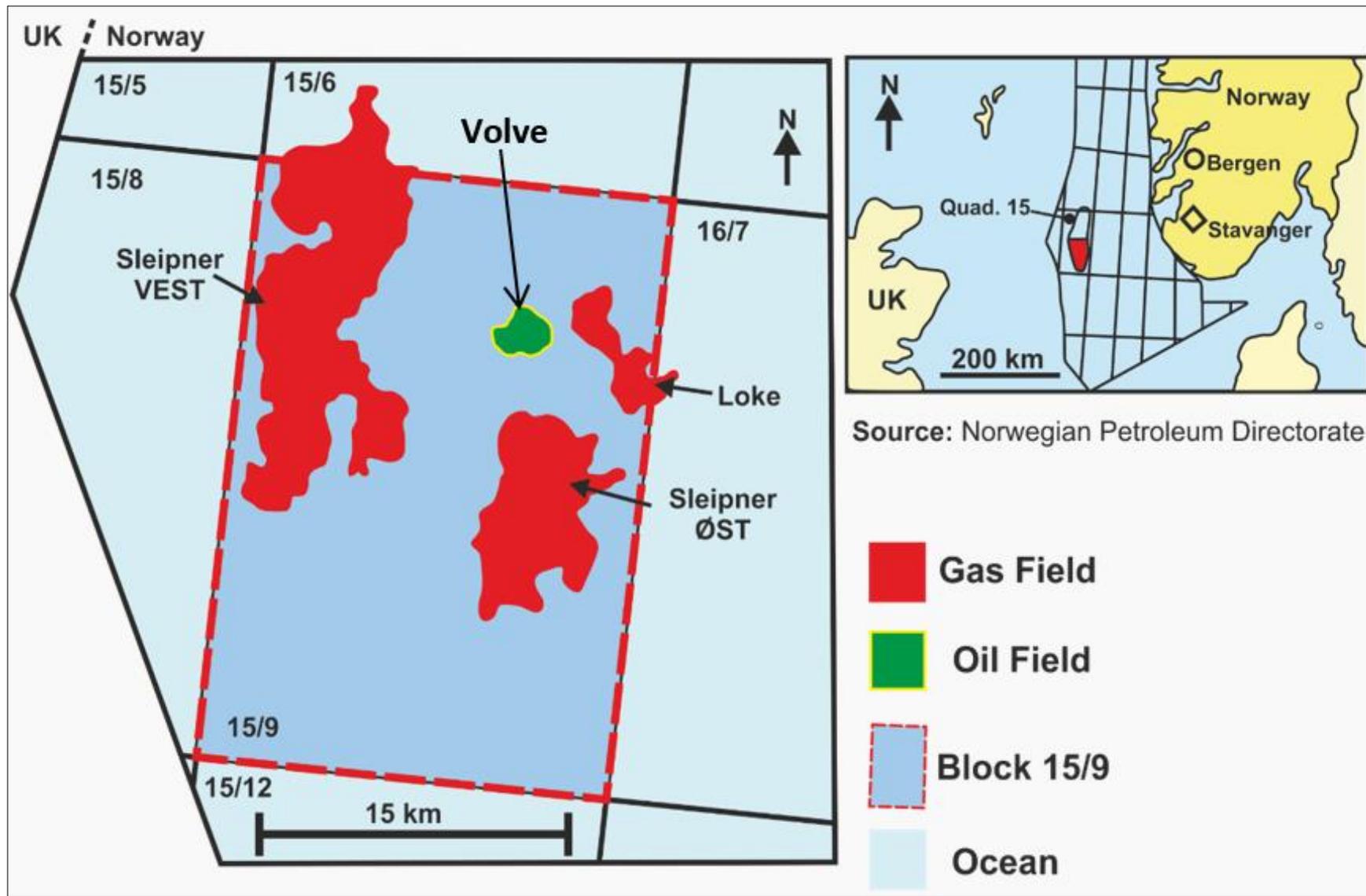
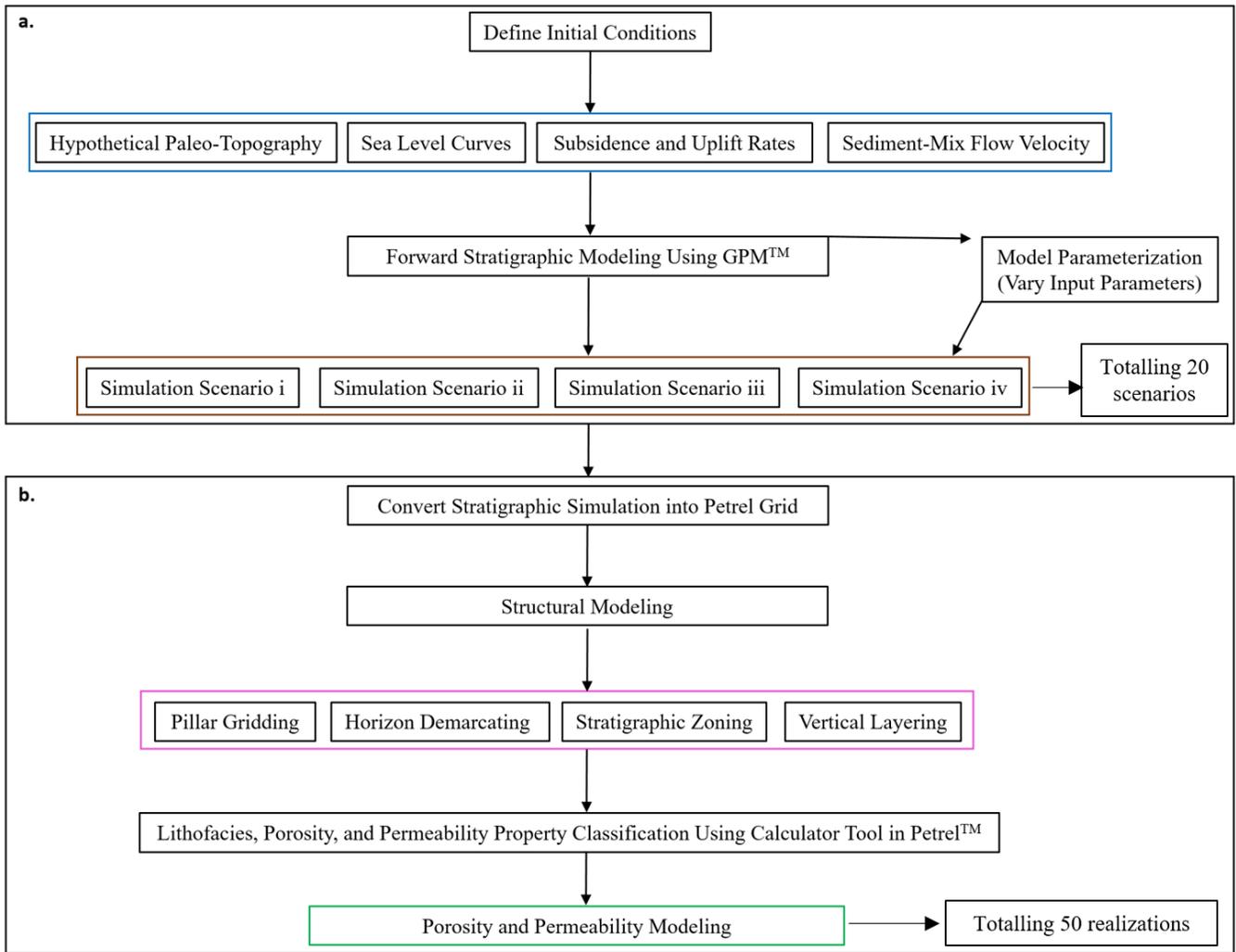


Fig 1. Location map of the Volve field, showing gas and oil fields in quadrant 15/9, Norwegian North Sea (Adapted from Ravasi et al., 2015).



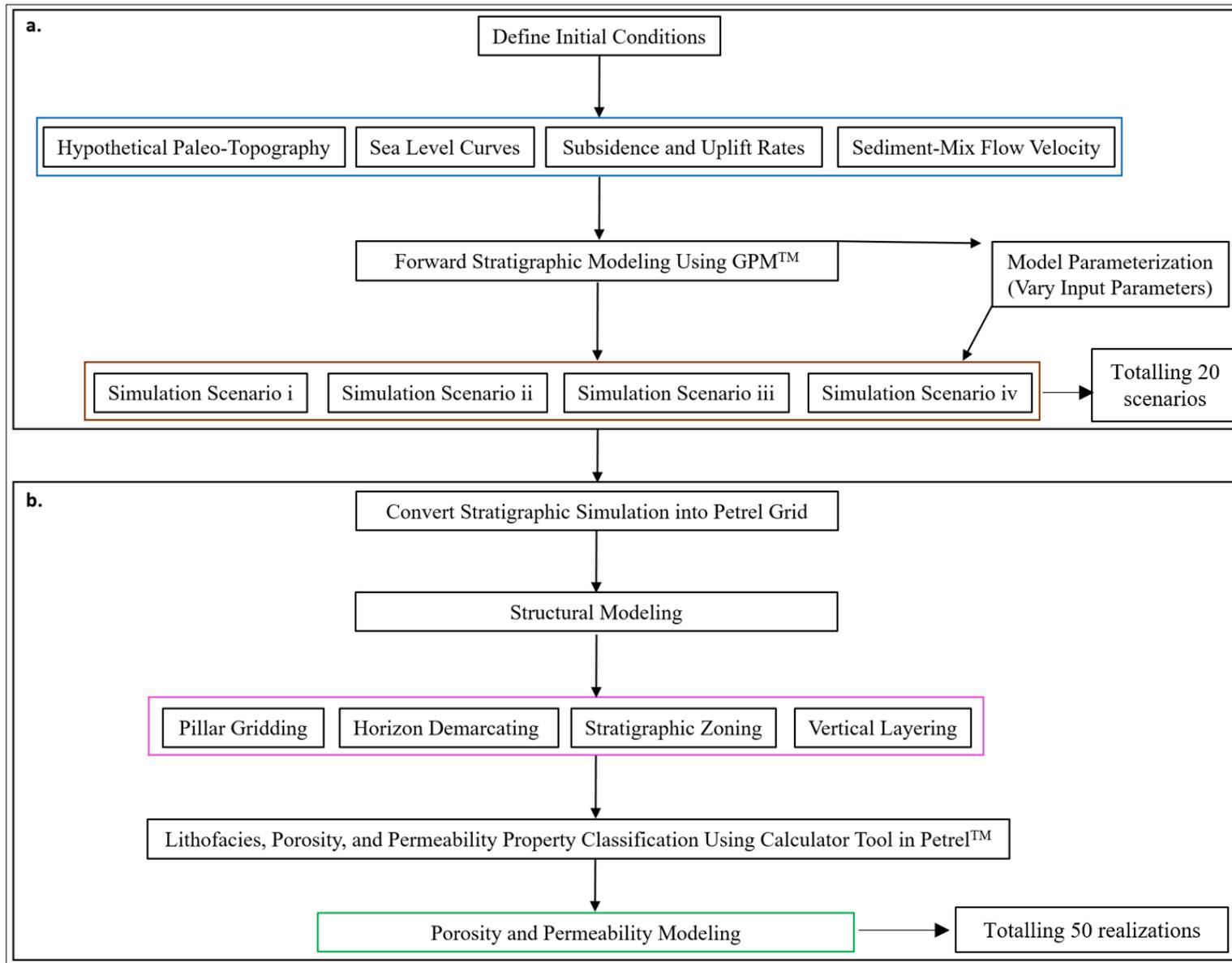
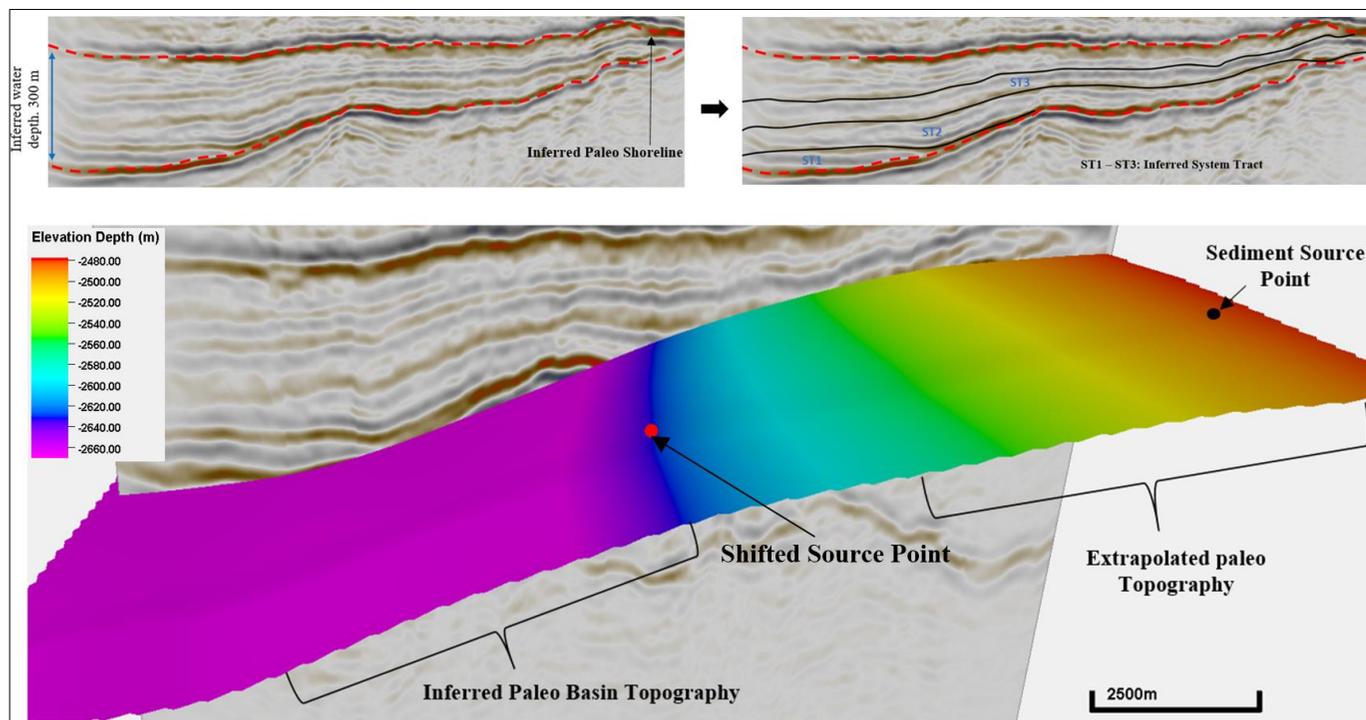


Fig 2. Schematic workflow of processes involved this work. a. providing information of initial conditions (or input parameters) that were used in the forward stratigraphic simulation in GPM™, b. demonstrating how the forward stratigraphic were converted into a grid that is usable in ~~the~~ Petrel™ ~~environment~~ for ~~onward 3-D~~ porosity and permeability modeling.



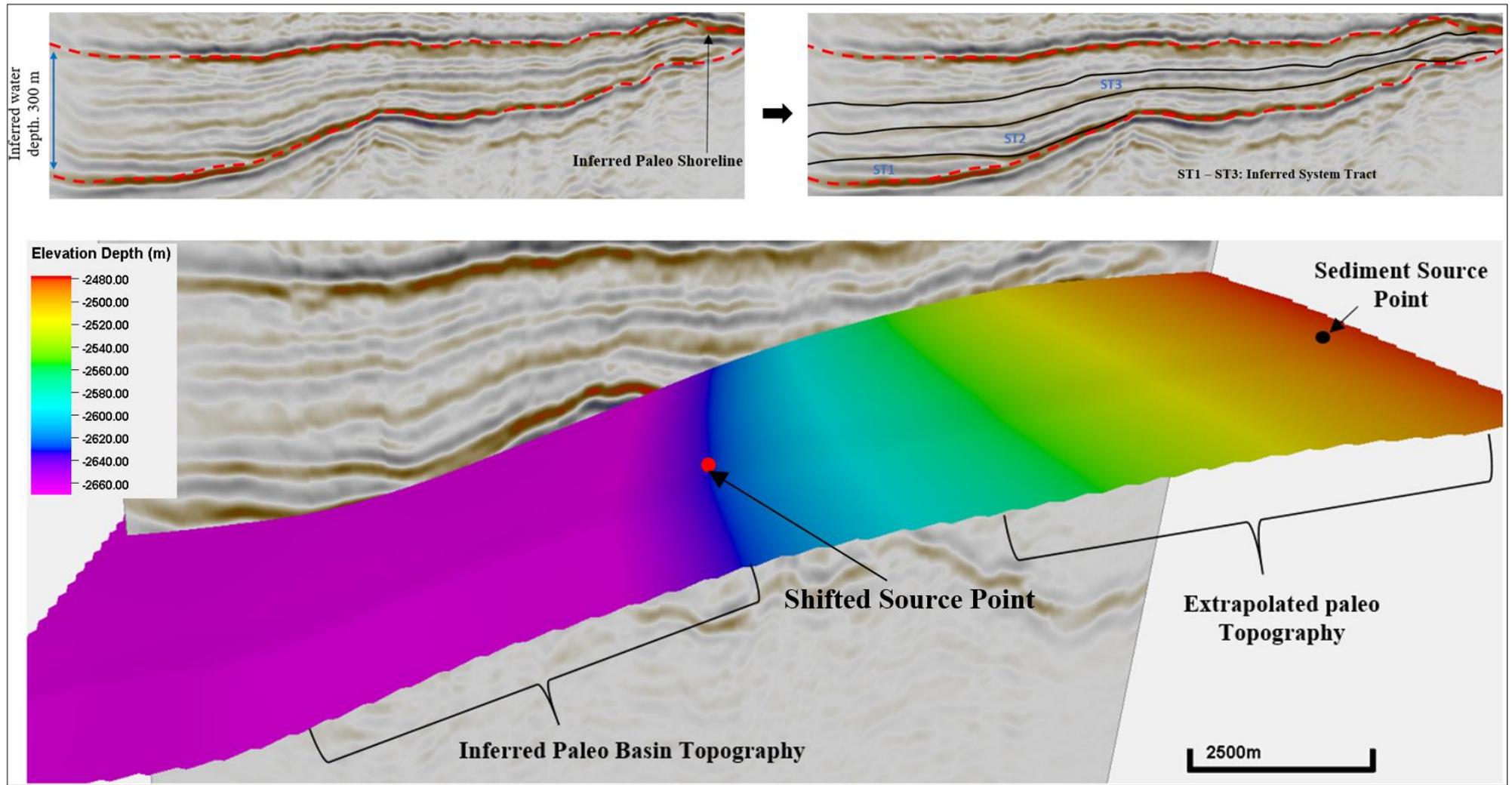
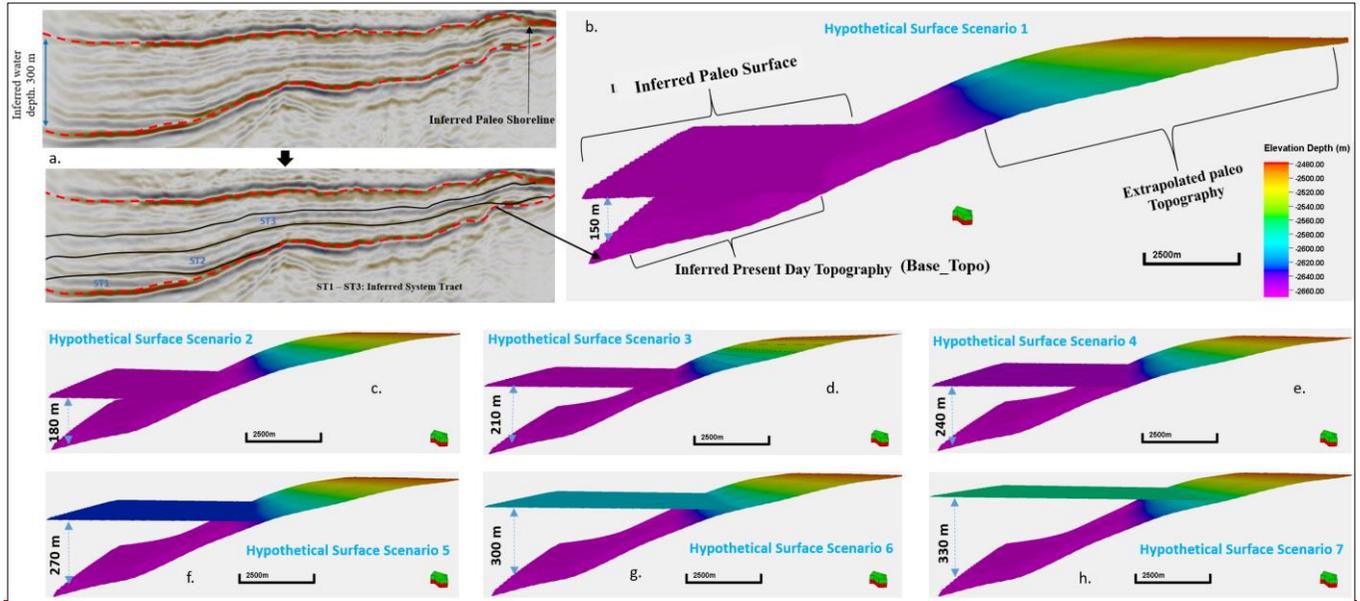


Fig 3. 3-D seismic section of the study area, from which the hypothetical topographic surface was derived for the simulation. The sedimentary entry point into the basin is located in the North Eastern section, based on previous study in the model area (e.g. Kieft et al. 2011).



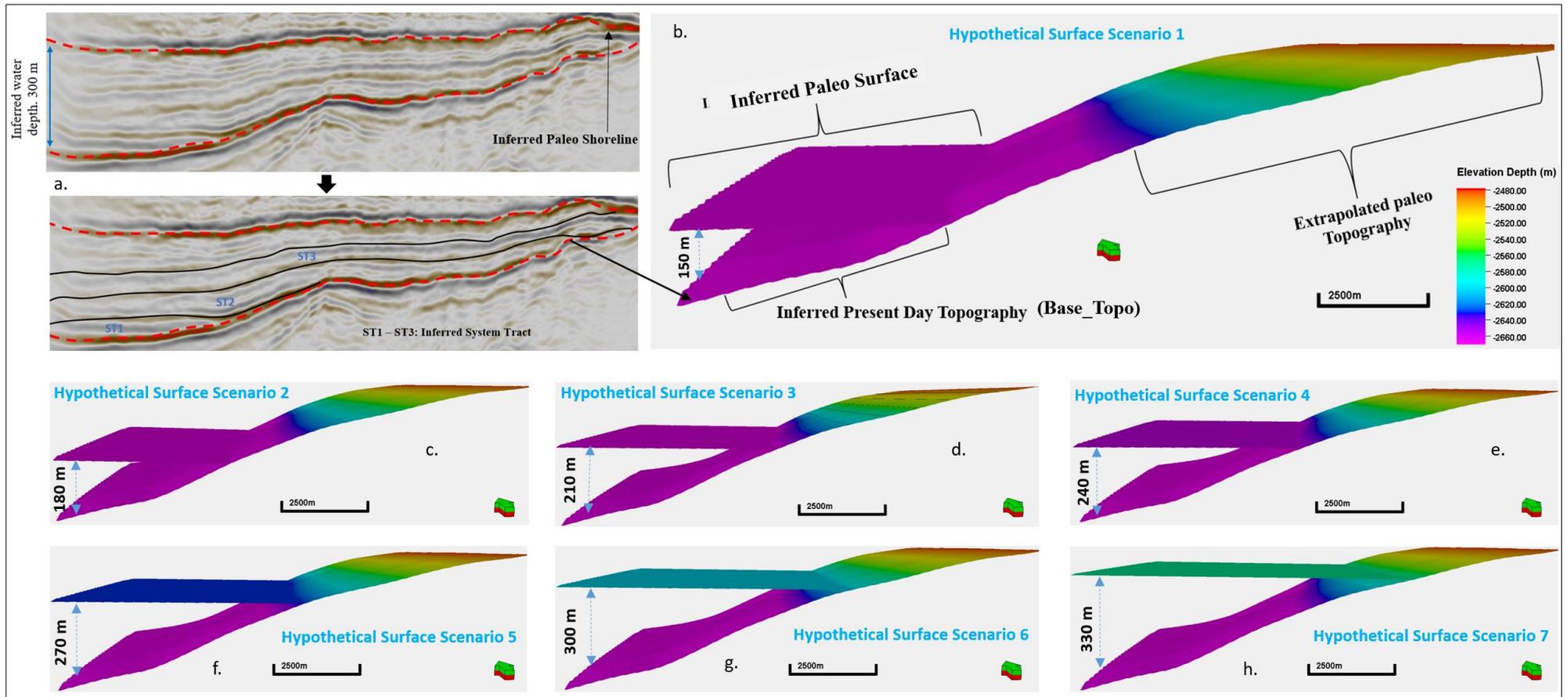
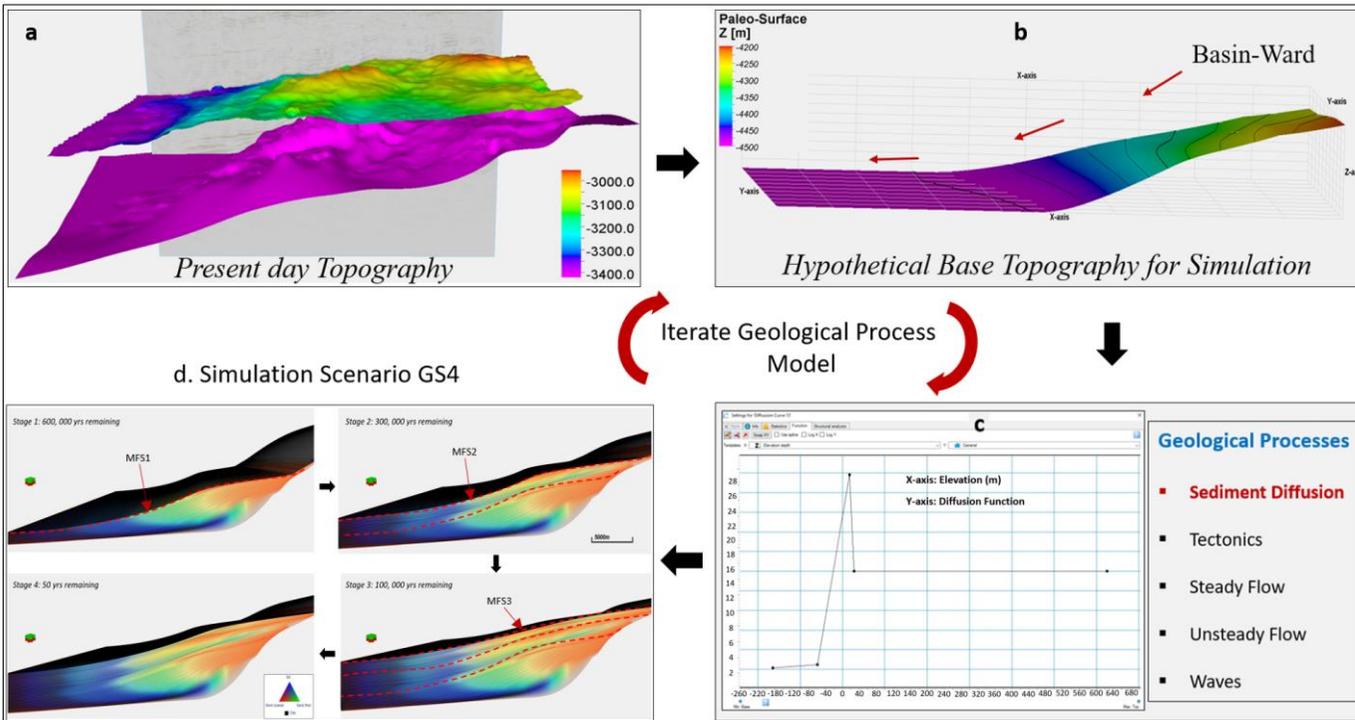


Fig 4. Inferred paleo topographic surface from seismic, also illustrating different topographic surface scenarios used in the simulation.



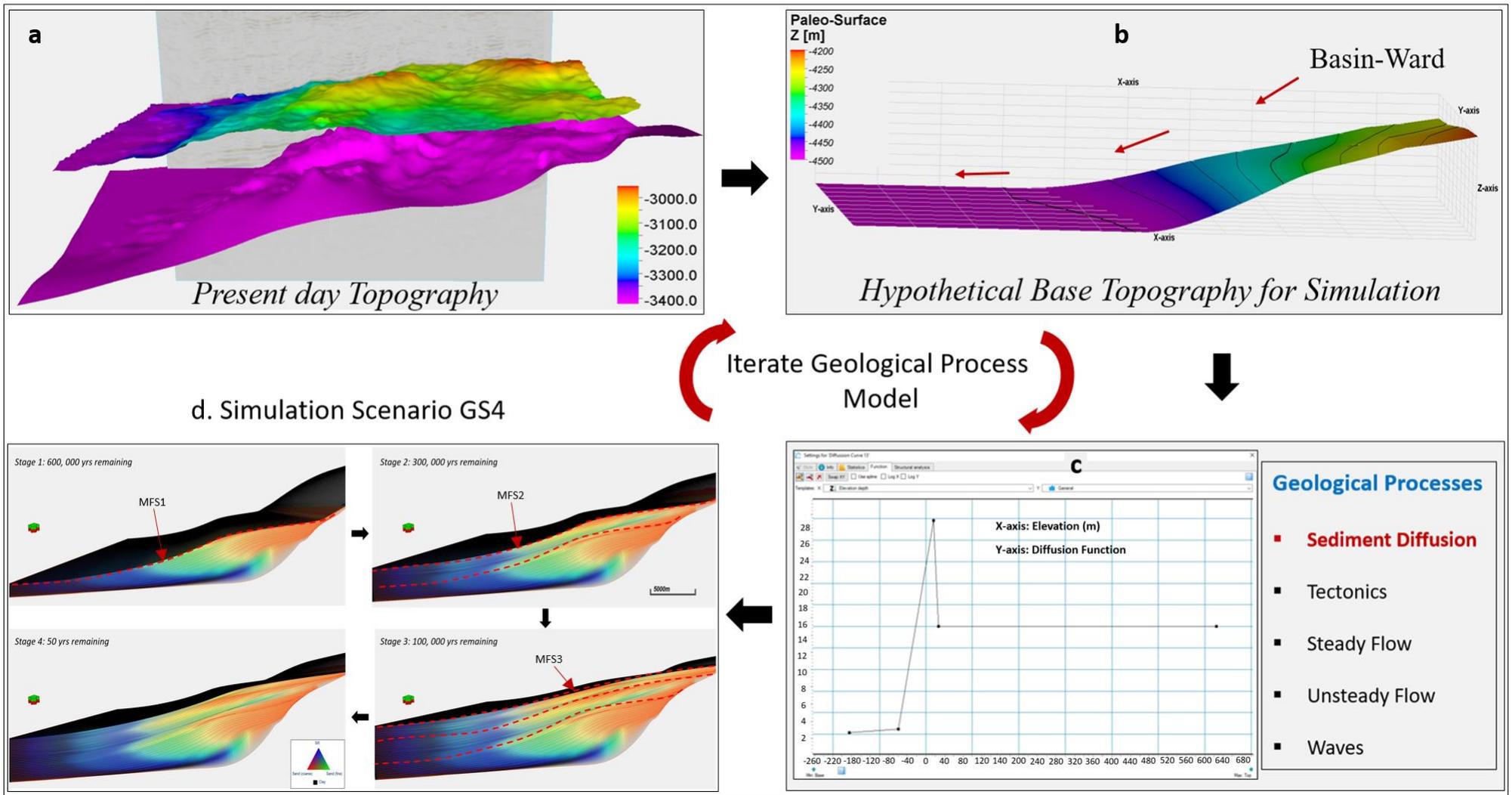


Fig 5. a. present day top and bottom topographic surfaces of the Hugin formation; b. hypothetical topographic surface derived from seismic data; c. geological processes involved in the simulation; d. forward stratigraphic models at different simulation time.

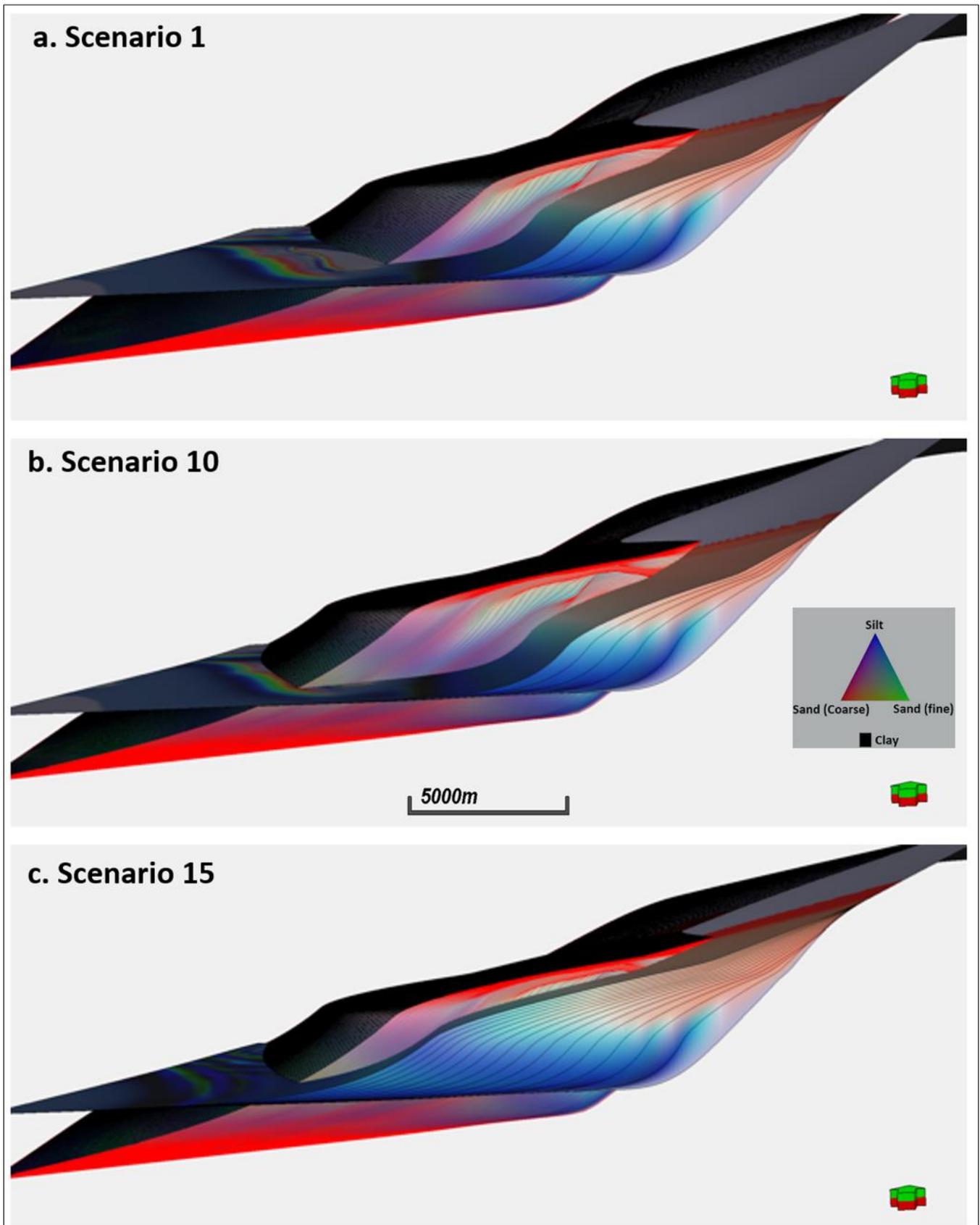
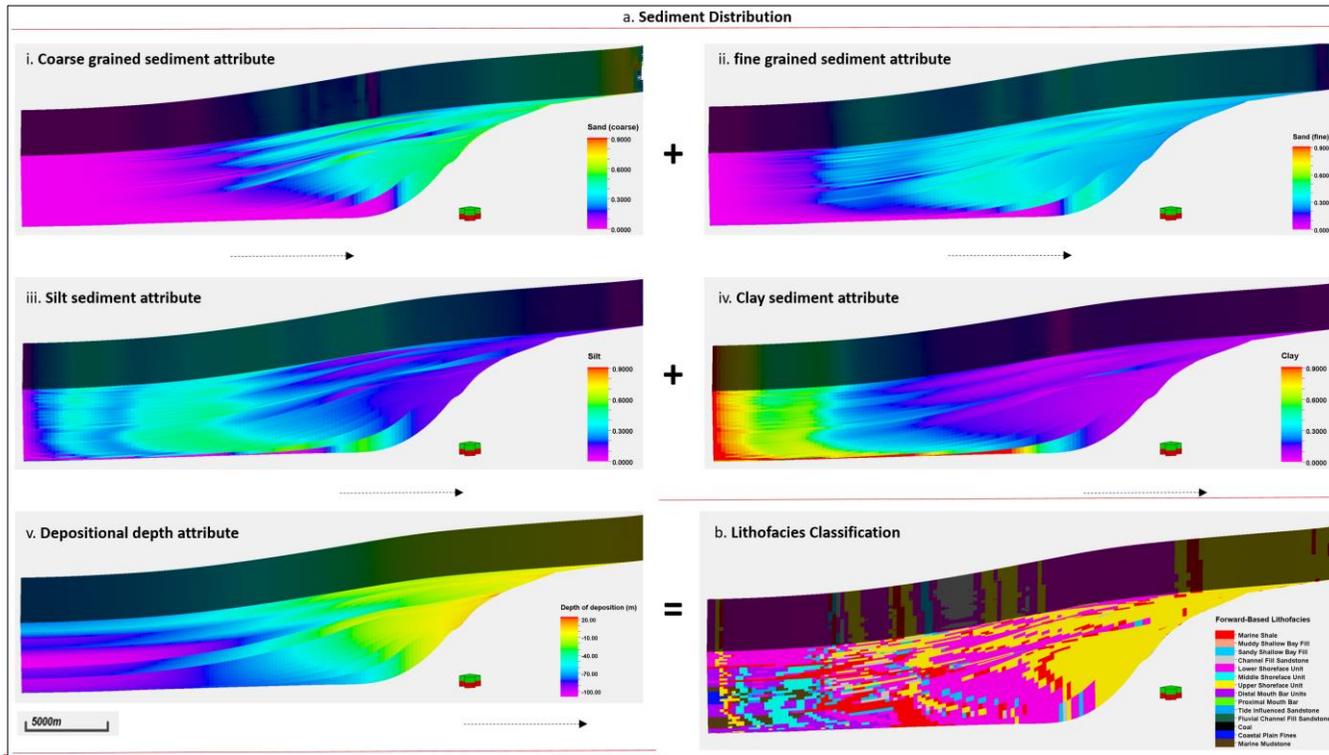


Fig 6. Stratigraphic simulation scenarios depicting sediment deposition in a shallow marine framework. **a.** scenario 1 involves equal proportions of sediment input, a relatively low subsidence rate and low water depth, **b.** scenario 10 uses high proportions of fine sand and silt (i.e. 70%) in the sediment mix, abrupt changes in subsidence rate, and a relatively high water depth, **c.** scenario 15 involves very high proportions of fine sand and silt (i.e. 80%), steady rate of subsidence and uplift in the sediment source area, and a relatively low water depth.



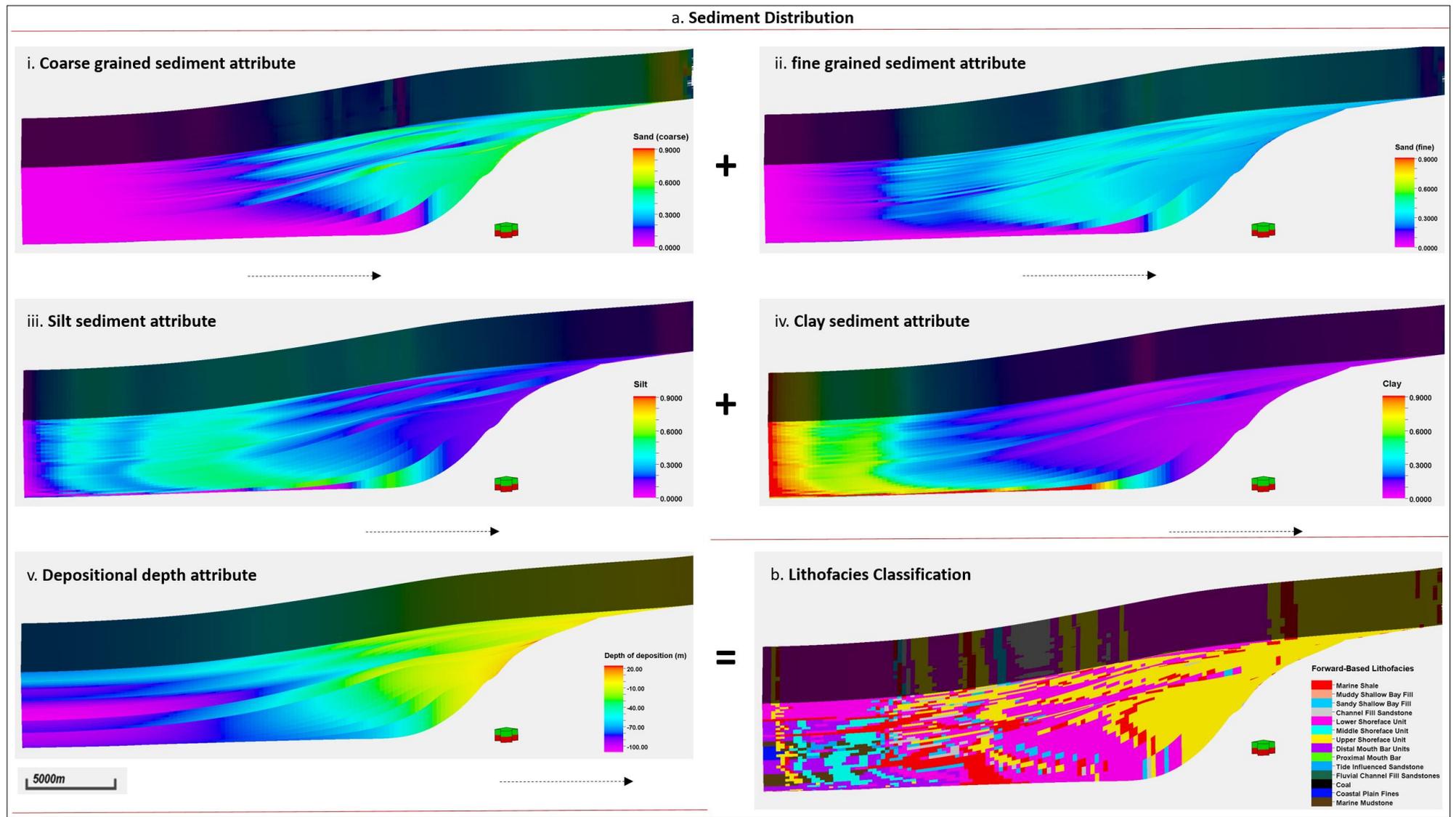


Fig 7 a. Sediment distribution patterns in the geological process modeling software. **b.** lithofacies classification using the property calculator tool in Petrel™.

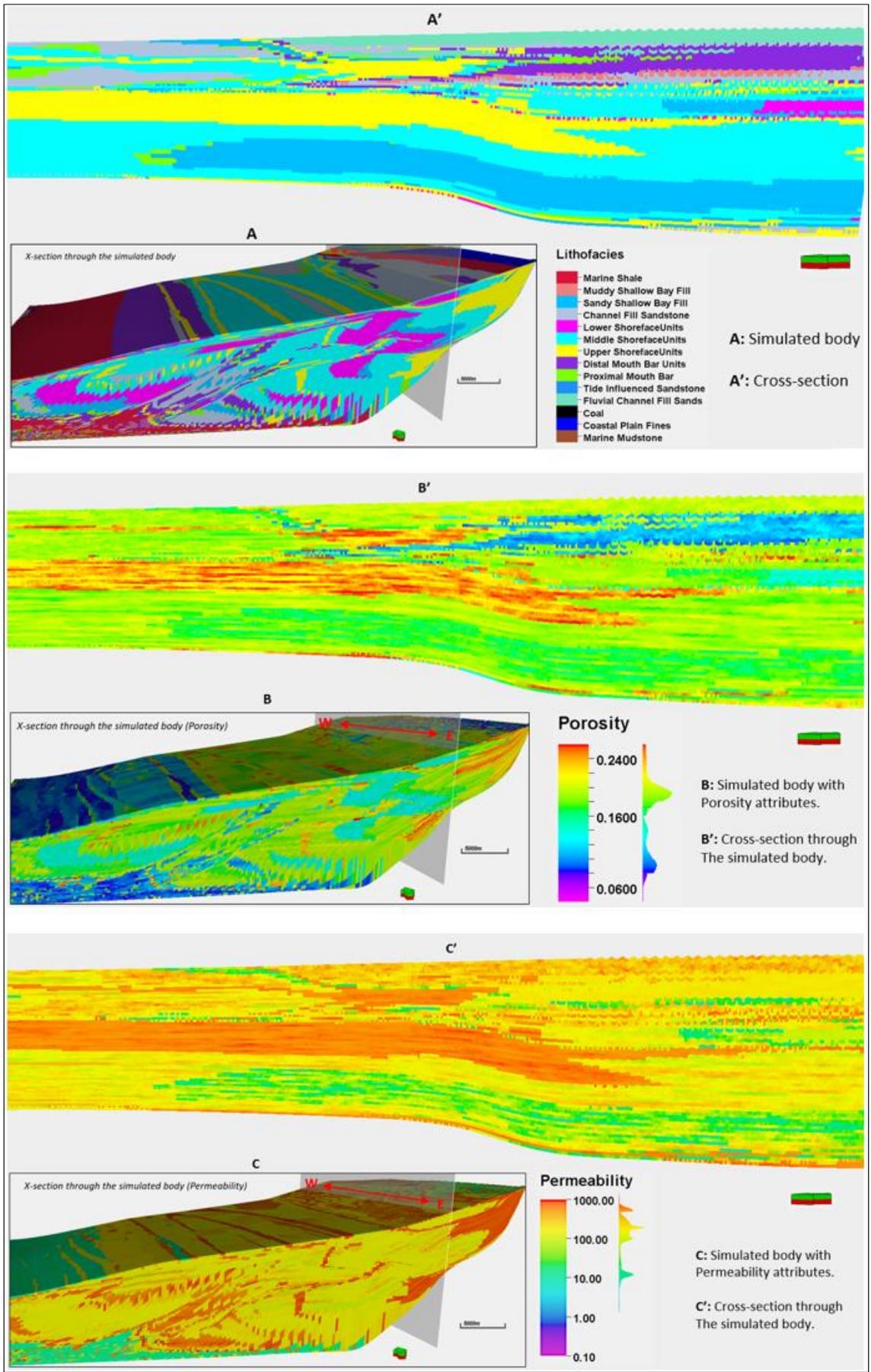
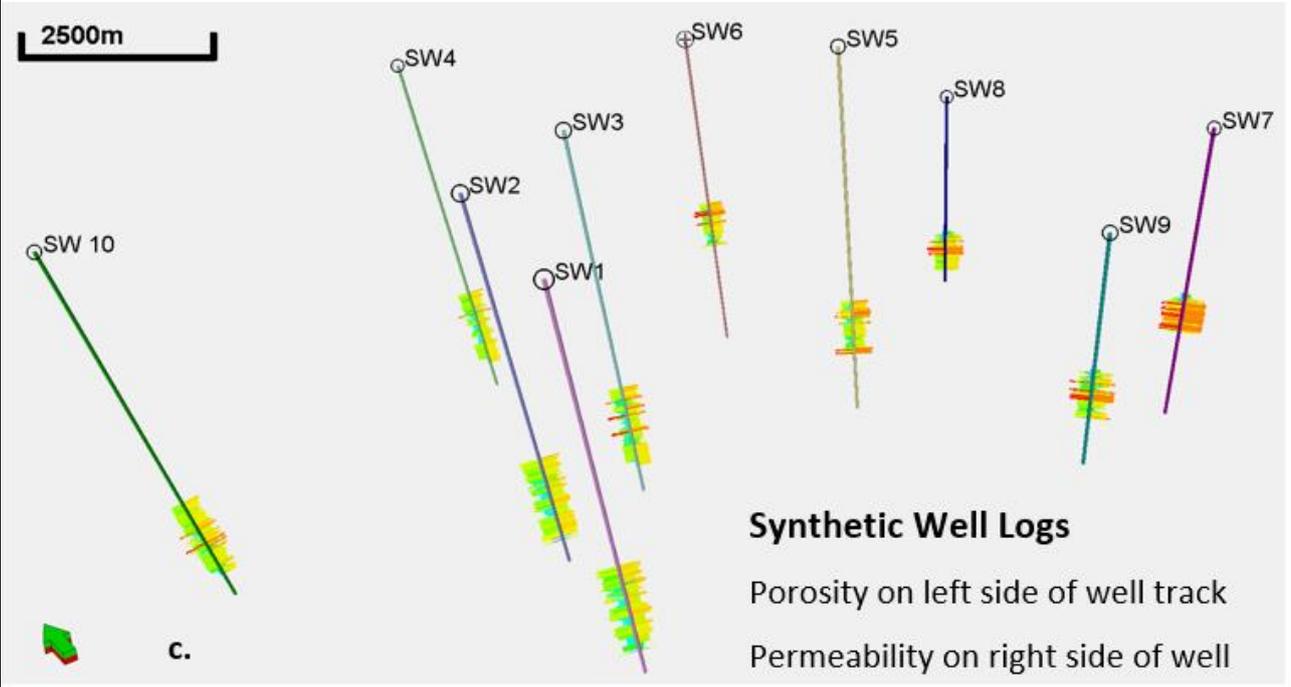
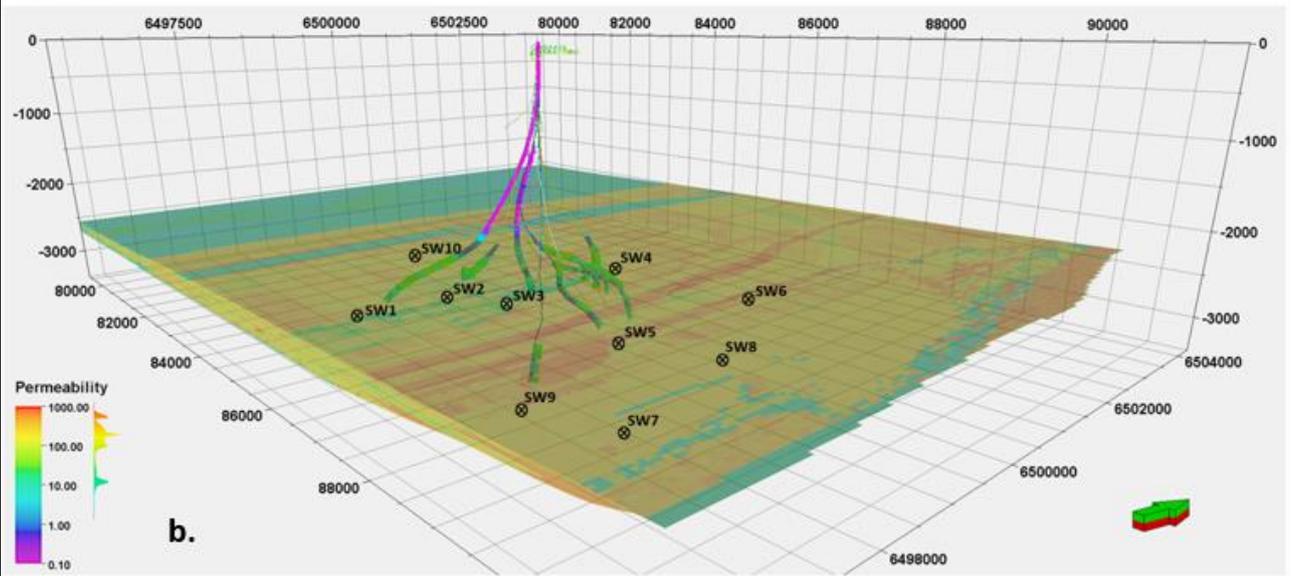
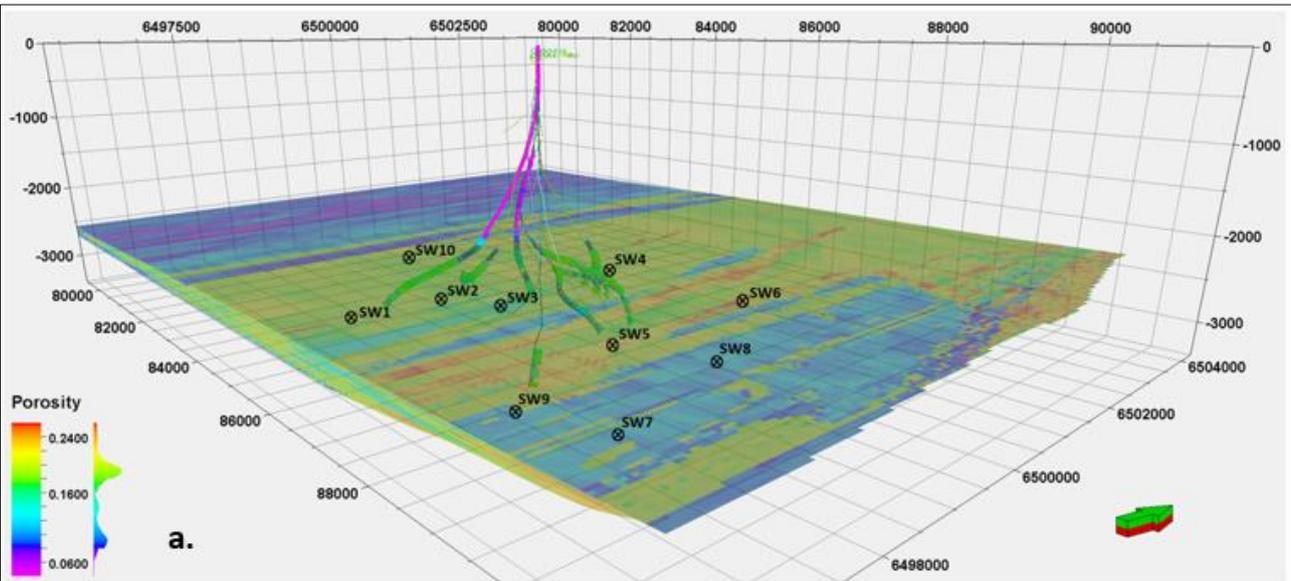


Fig 8. Property characterization in the stratigraphic using the property calculator tool in Petrel. Also showing a cross-sectional view through the model.



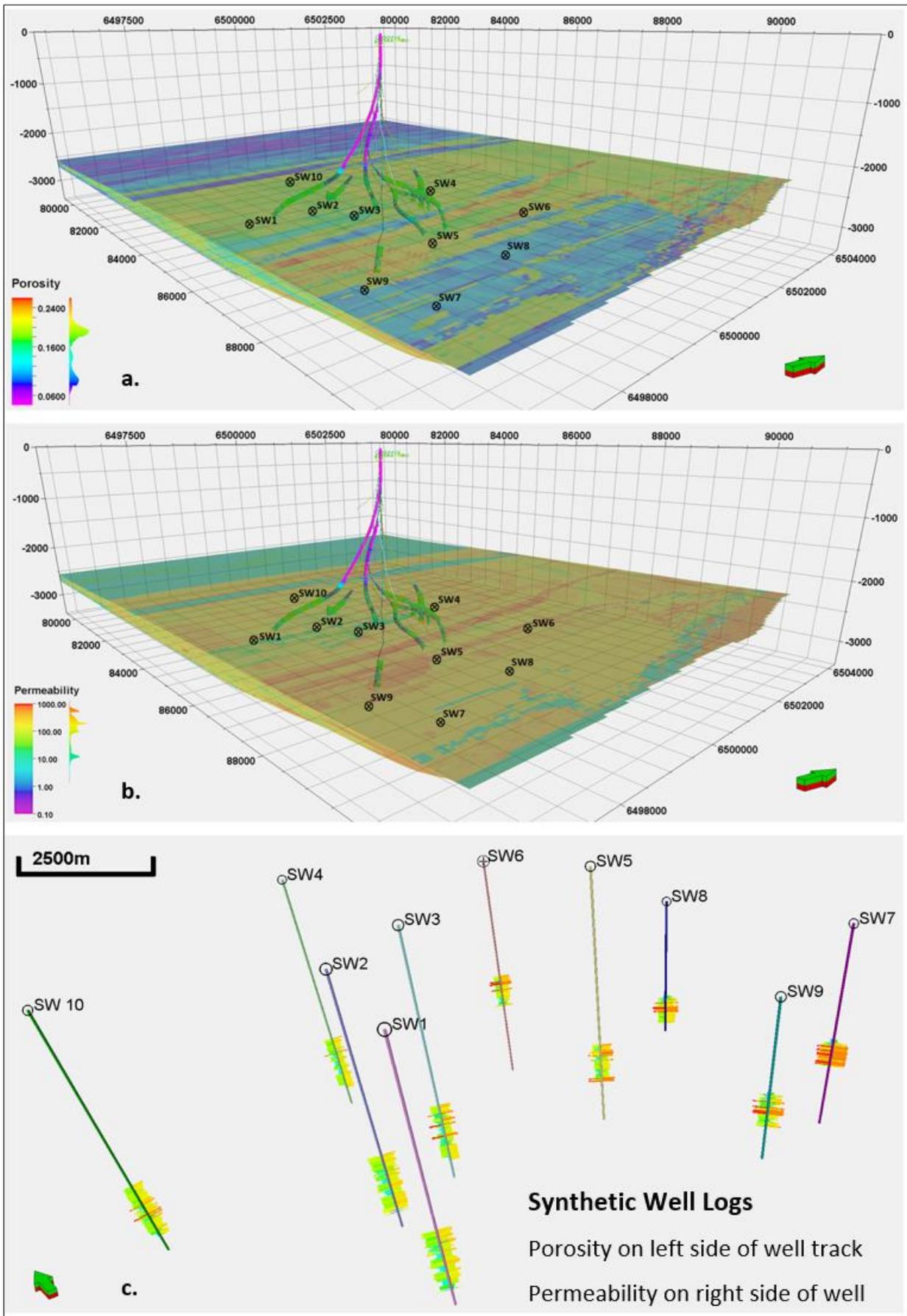
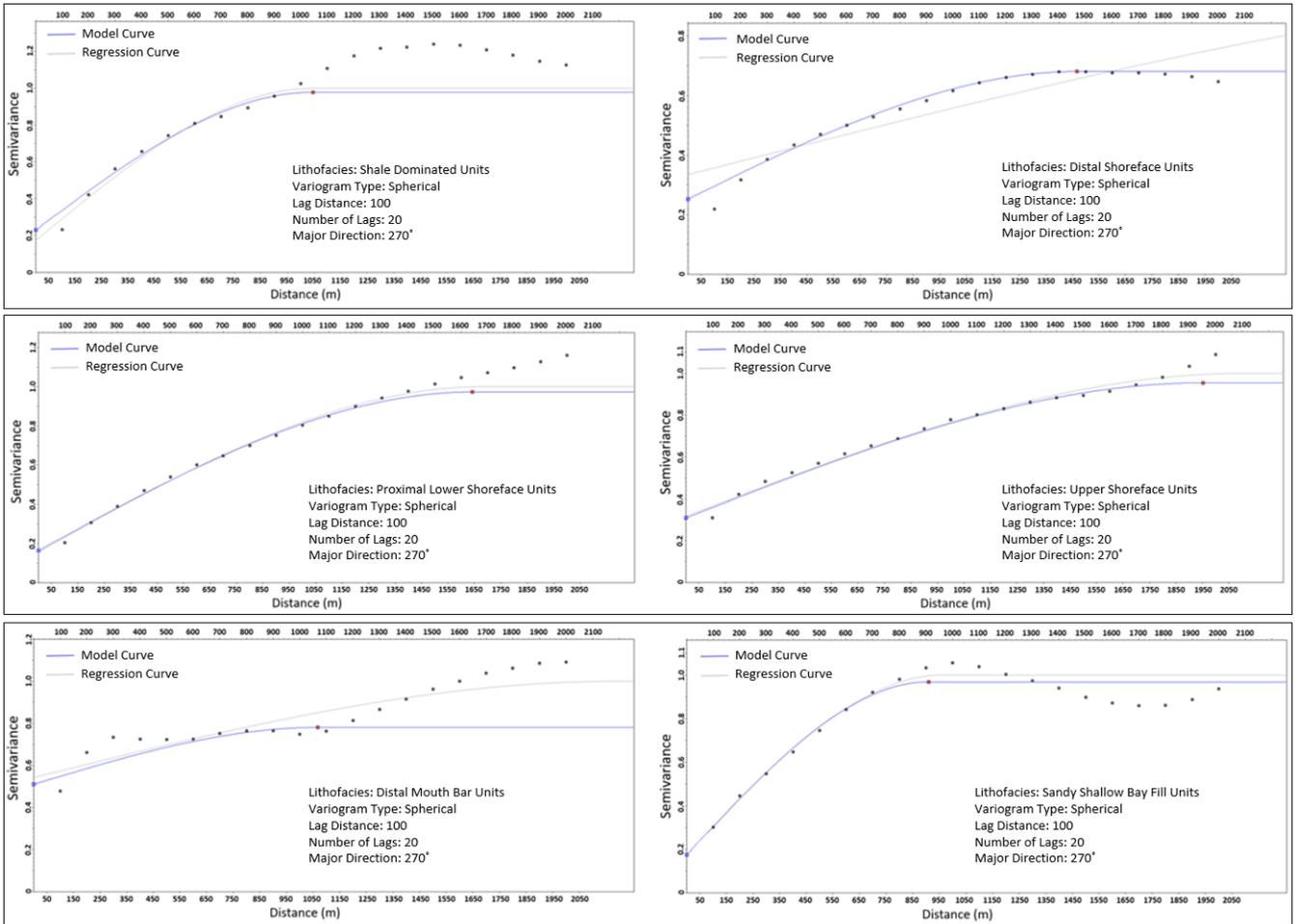


Fig 9. Synthetic wells derived from a forward stratigraphic-driven porosity and permeability models.



The average distance between the synthetic wells shown in Figure 9c is about 0.9 km apart (maximum and minimum separation distance of 1.3 km and 0.65 km respectively).

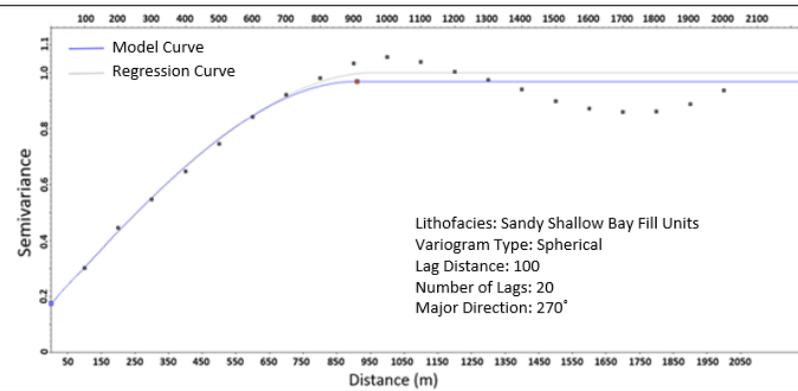
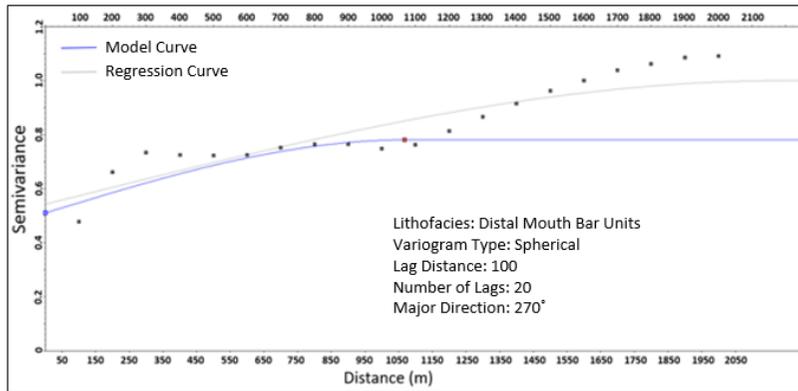
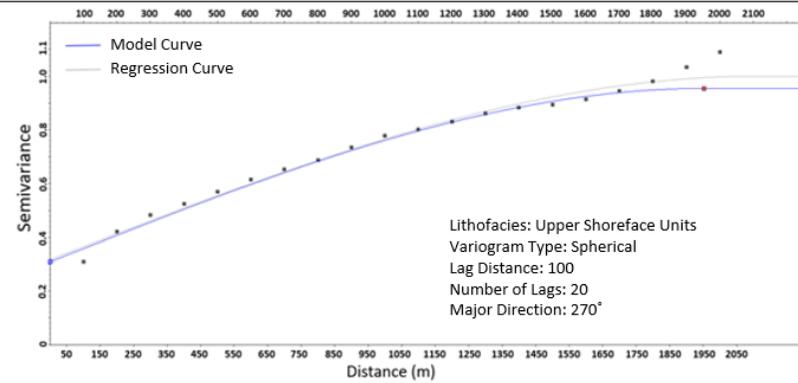
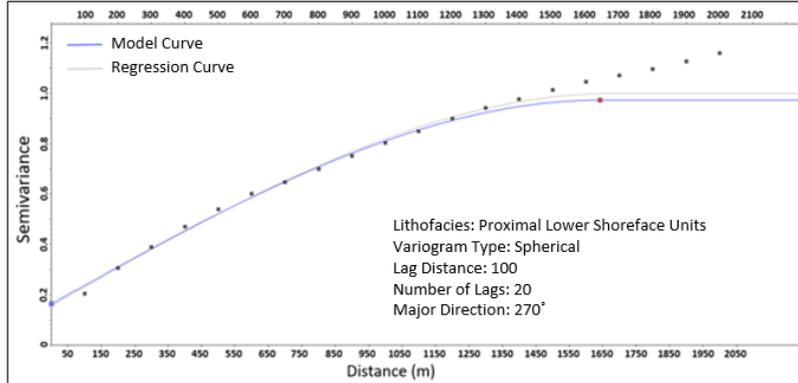
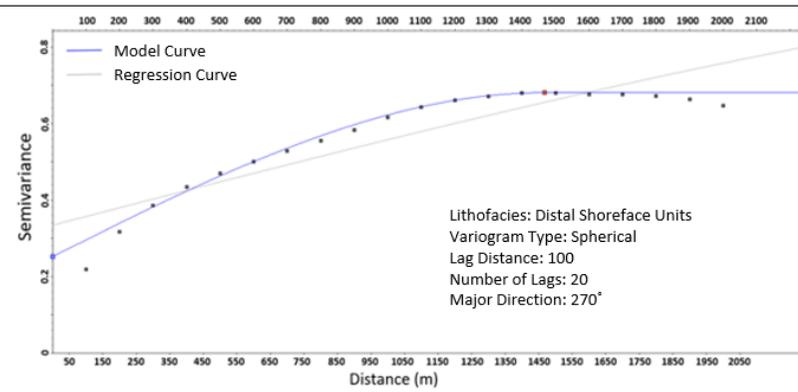
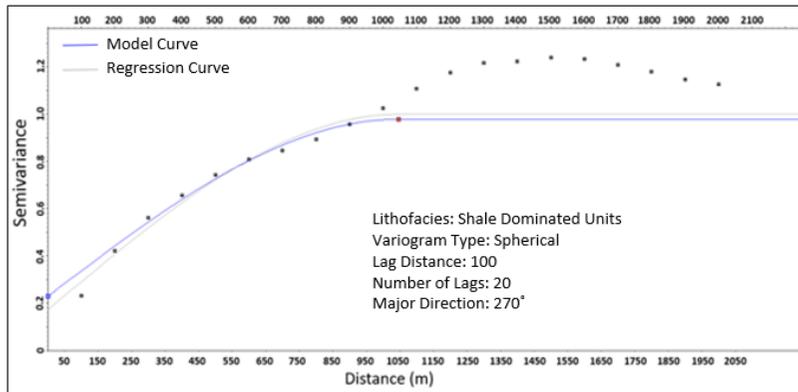
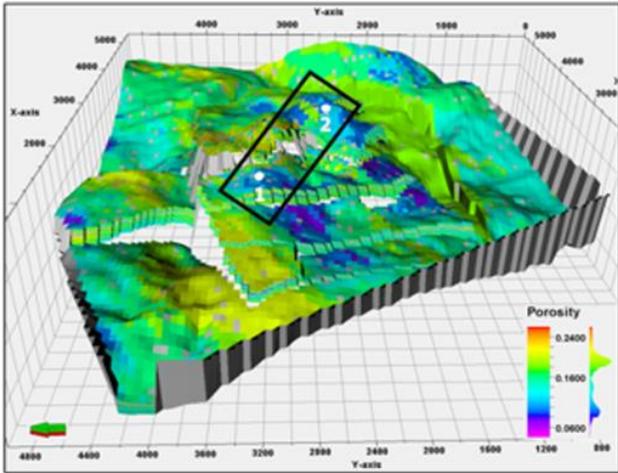
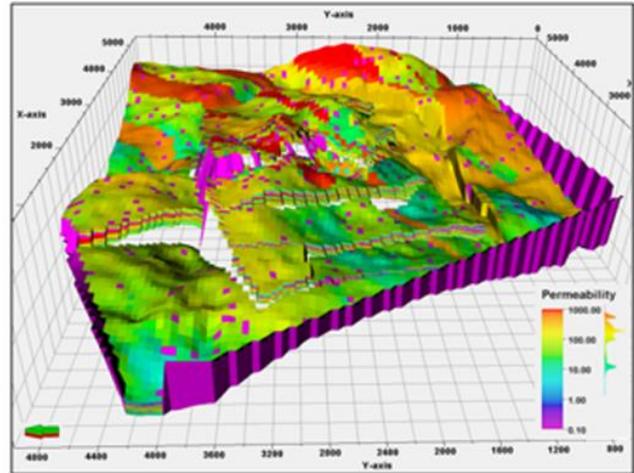


Fig 10. Variogram model of dominant lithofacies units extracted from the FSM. The points indicate the number of lags in the variogram. The distance between these lags is about 100 m. This figure shows the lags between sample pairs for calculating the variogram in the major direction (NE-SW) of the stratigraphic model.

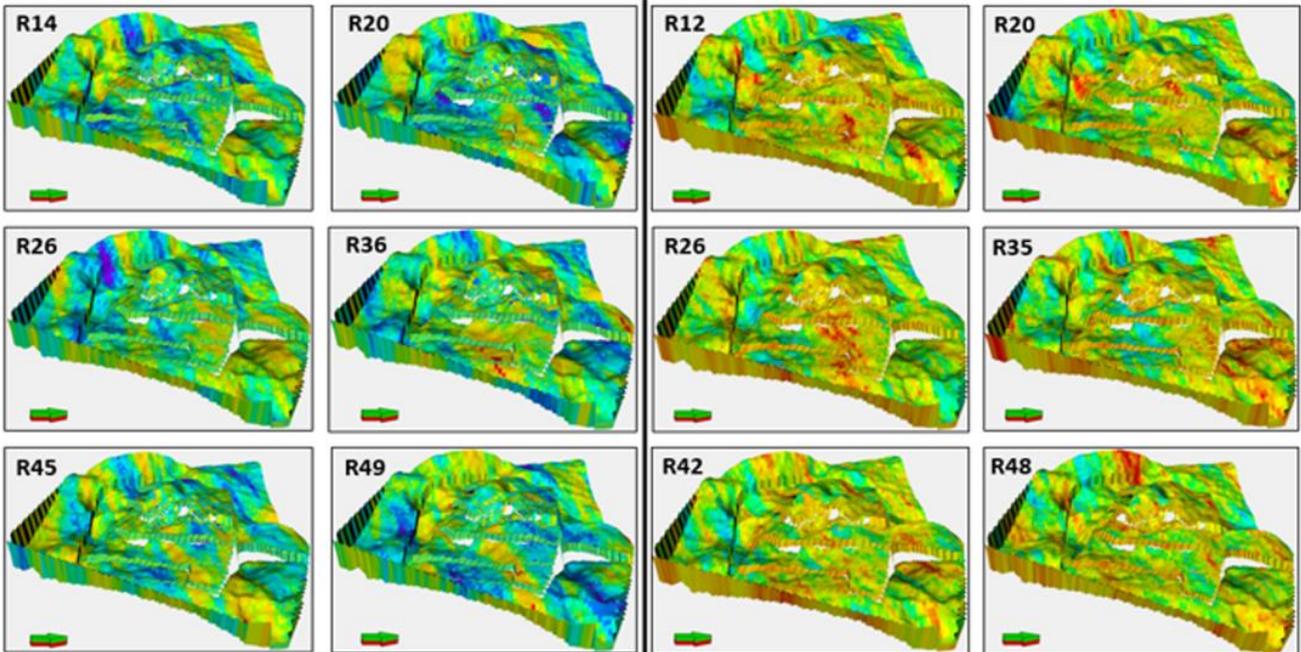
a. Original Porosity Model



Original PermZ Model



b. Forward Modeling-Based Porosity and Permeability Model Realizations



c.

Validation Wells Location

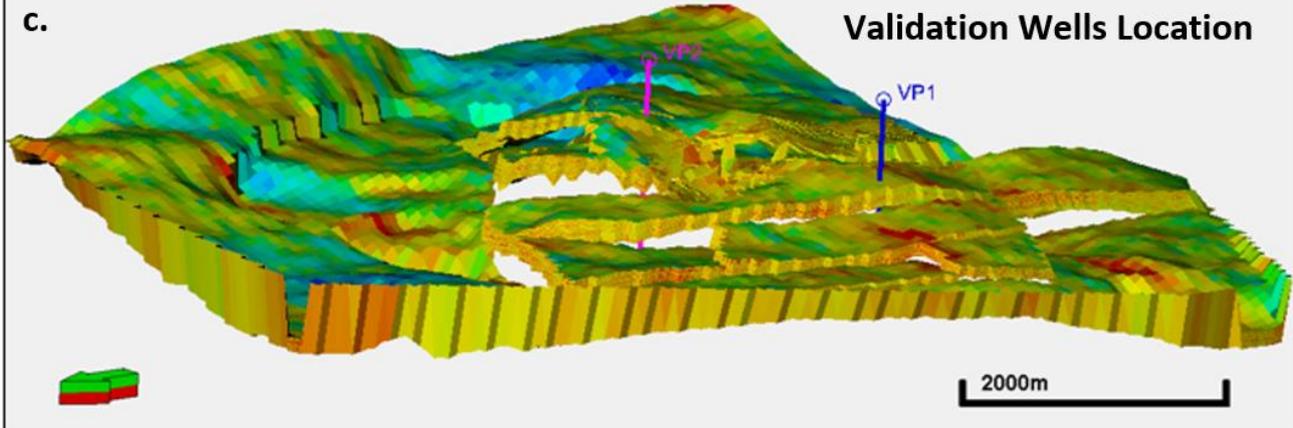
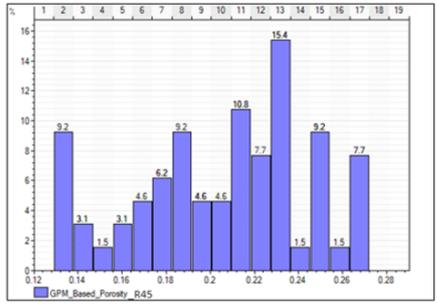
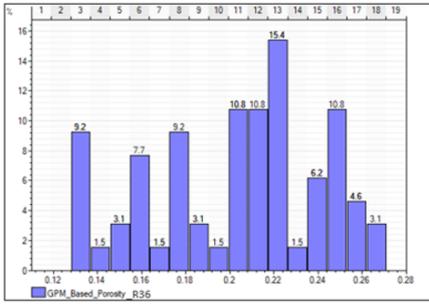
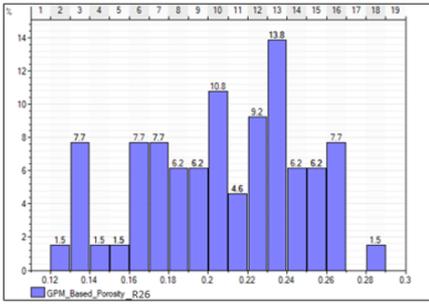
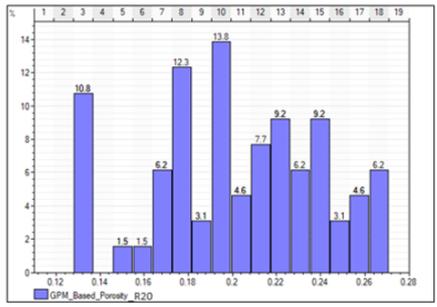
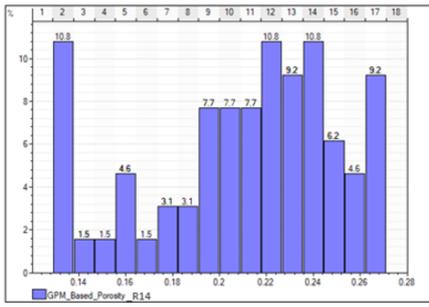
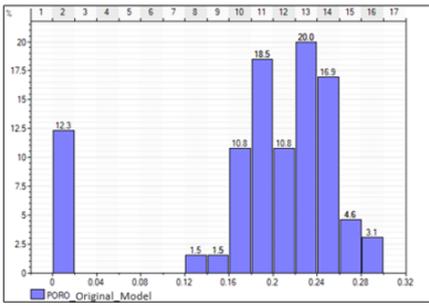
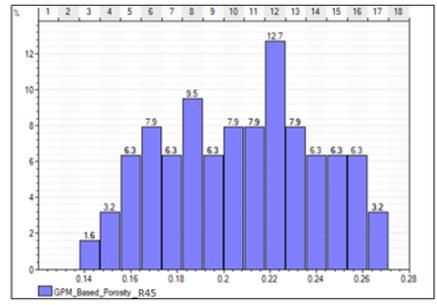
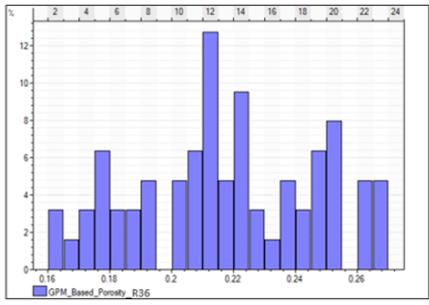
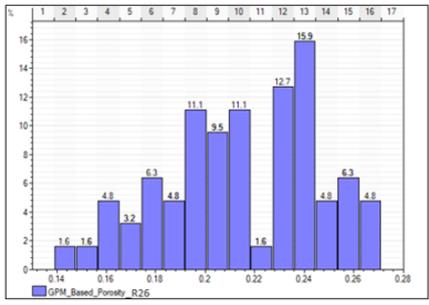
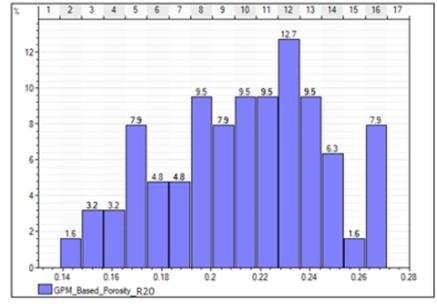
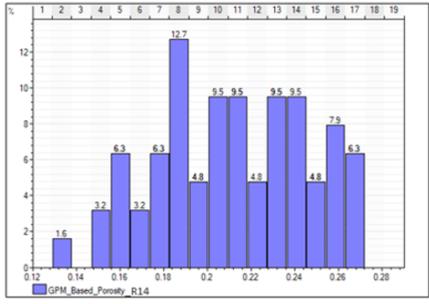
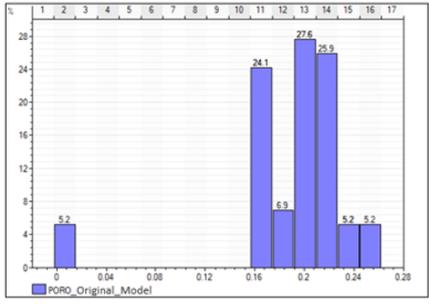


Fig 11. Comparing original Volve field model to the forward modeling-based models. Realizations 16, 20, 26, 36, 45, and 49 on the left half are porosity models, whilst realizations 12, 20, 26, 35, 42, and 48 on the right half show permeability models.

a. Validation Well 1



b. Validation Well 2



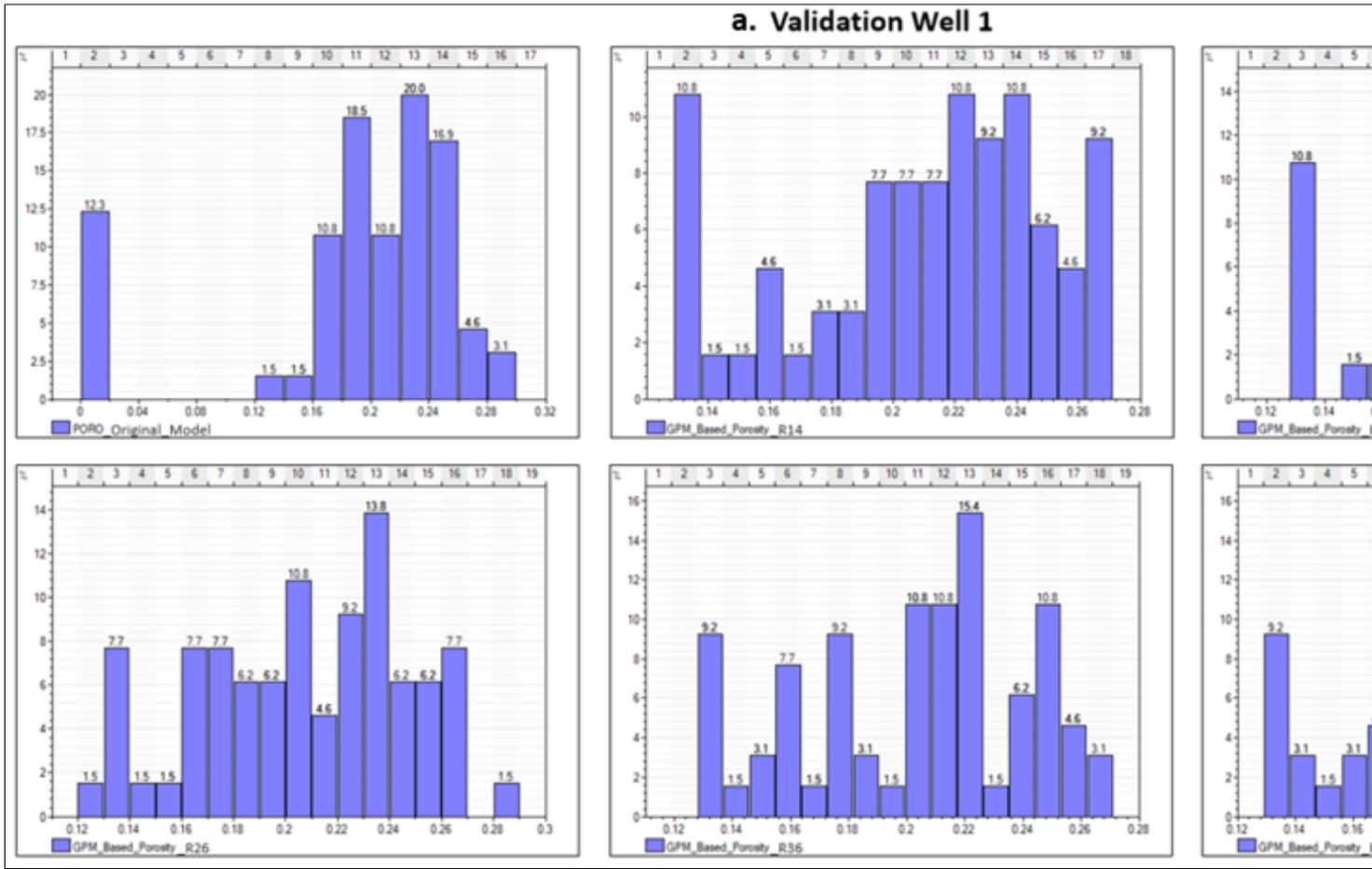


Figure 12. Illustrating how; a. Validation well 1 in five selected realizations, and b. validation well 2 samples in the synthetic forward-based model how it compares to pseudo-wells from the samples at similar vertical interval in the original Volve field petrophysical model.

porosity and permeability models.

b. Validation Well 2

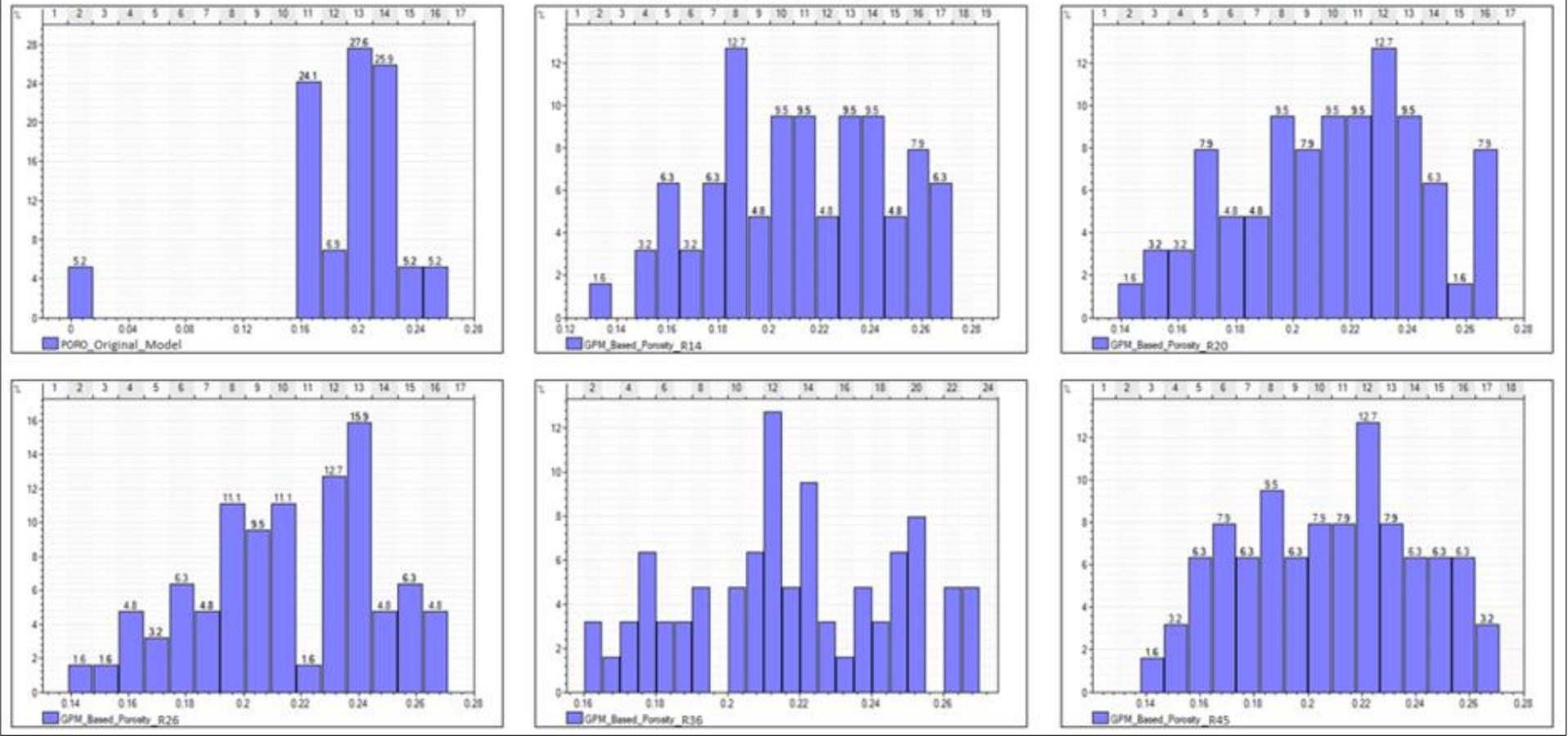


Figure 12b. Samples of validation Well 2 in five selected realizations, and how it compares to the samples at similar vertical interval in the original porosity and permeability models.

Table 1 Lithofacies-associations in the Hugin formation, Volve Field (after Kieft et al. 2011).

Code	Facies	Description	Thickness (t); extent (l)	Wireline-log Attribute	Interpretation
A	A1	Parallel-laminated mudstone with occasional siltstone inputs. Monospecific pattern of disorder bivalves parallel to bedding.	t= 30-425 cm l= 6 to 29 km	GR= 41-308 API DT= 225-355 μsm^{-1} NPHI= 0.17-0.45 v/v RHOB= 2280-2820 gcm^{-3}	Restricted marine shale
	A2	Interbedded claystone and very fine-grained sandstone; non-parallel and wavy lamination. Scarcely bivalve shells oriented parallel to bedding.	t= 10-725 cm l= 8 km to 13 km	GR= 71-65 API DT= 189-268 μsm^{-1} NPHI=? RHOB= 2280-2820 gcm^{-3}	Muddy Shallow bay-fill
	A3	Fine to medium grained sandstone; moderately to well sorted grains. Wavy bedding, cross bedding, rare wave ripples	t= 60-370 cm l < 8 km	GR= 18-46 API DT= 199-314 μsm^{-1} NPHI= 0.07-0.52 v/v RHOB= 1690-2745 gcm^{-3}	Sandy shallow bay-fill
	A4	Coarse to fine-grained sandstones with alternating upward fining to coarsening trend. Moderately sorted grains. Sparse sedimentary structures.	t= 250-500 cm l= 1.8 km to 4.2 km	GR= 7-35 API DT= 175-230 μsm^{-1} NPHI= 0.038-0.146 v/v RHOB= 2280-2820 gcm^{-3}	Marine channel-fill sandstones
B	B1	Upward-coarsening siltstone to fine-grained moderate sorted sandstones, with shell debris, and quartz granules.	t= 30-480 cm l = <2 km	GR= 18-80 API DT= 168-291 μsm^{-1} NPHI= 0.038-0.191 v/v RHOB= 2322-2723 gcm^{-3}	Distal lower shoreface
	B2	Very fine-grained, moderate to well sorted sandstone. Fine grained carbonaceous laminae, typically low angle cross beds.	t= 130-440 cm l = 1.7 km – 8 km	GR= 20-56 API DT= 179-277 μsm^{-1} NPHI= 0.048-0.168 v/v RHOB= 2314-2696 gcm^{-3}	Proximal lower shoreface
	B3	Coarsening upward, cross laminated, fine to medium grained, well sorted sandstone; consist carbonaceous fragments	t= 425-800 cm l = 1.7 km – 8 km	GR= 15-25 API DT= 250-275 μsm^{-1} NPHI= 0.09-0.113 v/v RHOB= 2271-2342 gcm^{-3}	Upper Shoreface
C	C1	Highly bioturbated siltstone to very fine sandstones, which has beds of rounded granules	t= 175-1010 cm l = 7.2 km – 19.6 km	GR= 20-80 API DT= 230-260 μsm^{-1} NPHI= 0.08-0.169 v/v RHOB= 2327-2521 gcm^{-3}	Distal mouth bar
	C2	Very fine to fine grained sandstones; low angle cross-bedding.	t= 290-775 cm l = < 5 km	GR= 12-58 API DT= 167-397 μsm^{-1} NPHI= 0.05-0.595 v/v RHOB= 1612-2705 gcm^{-3}	Proximal mouth bar
D	D1	Fining upward coarse to fine grained sandstone; stacked fining upward beds with rare coarse grained stringers.	t= 740-820 cm l = < 2 km	GR= 8-134 API DT= 235-335 μsm^{-1} NPHI= 0.14-0.460 v/v RHOB= 2284-2570 gcm^{-3}	Tidally influenced fluvial channel fill sandstone
	D2	Fining upward coarse to medium grained sandstone. Carbonaceous laminae and fragments. Sharp, and cohesive contact at bed base	t= 580 cm l = < 2 km	GR= 9-34 API DT= 241-297 μsm^{-1} NPHI= 0.14-0.289 v/v RHOB= 2168-2447 gcm^{-3}	fluvial channel fill sandstone
E	E1	Coal and carbonaceous shale. Basal contact, typically parallel.	t= 30-520 cm l = 6 km to 19.6 km	GR= 8-56 API DT= 313-427 μsm^{-1} NPHI= 0.24-0.529 v/v RHOB= 1930-2225 gcm^{-3}	coal
	E2	Alternating dark grey mud/claystone and siltstone to very fine-grained sandstone. Wavy to non-parallel lamination.	t= 60 cm l = < 2 km	GR= 32-60 API DT= 358-415 μsm^{-1} NPHI= 0.43-0.49 v/v RHOB= 1994-2148 gcm^{-3}	Coastal plain fines

Table 2. Input parameters applied in running the simulations in GPM™

Initial Conditions- GPM Input Parameters														
	Simulation Duration (Ma- 0a) Years	Sediment Type Proportion (%)				Avg. Water Velocity	Avg. Sediment Velocity	Erodibility	Diffusion Coefficient	Avg. Sea Level	Turbidite Event Interval	Steady Flow Iteration	Sediment Movement	
		Sand (Coarse)	Sand (Fine)	Silt	Clay	(m/a)	(m/a)			Interval (m)	(/years)	(/hrs)	Coefficient	
GPM Scenarios (GS)	S1	0.02-0	25	25	25	25	0.11	0.03	0.35	0.11	30	2500	10	0.001
	S2	0.25-0	25	25	25	25	0.15	0.03	0.45	0.15	70	1000	15	0.012
	S3	0.5-0	25	25	25	25	0.11	0.02	0.55	0.11	120	1000	20	0.012
	S4	0.7-0.05	25	25	25	25	0.08	0.02	0.35	0.08	100	500	25	0.0011
	S5	1.5-0	15	35	30	20	0.15	0.04	0.50	0.15	80	5000	20	0.001
	S6	3.0-0	50	25	15	10	0.13	0.04	0.50	0.13	70	5000	30	0.0012
	S7	3.5-0	50	25	15	10	0.11	0.04	0.50	0.11	70	10000	15	0.001
	S8	4.0-0	50	25	15	10	0.13	0.04	0.50	0.13	90	5000	20	0.0015
	S9	4.5-0	15	45	25	15	0.1	0.02	0.45	0.1	50	10000	30	0.0012
	S10	5.0-0	15	45	25	15	0.12	0.02	0.45	0.12	55	10000	35	0.0013
	S11	5.5-0	15	45	25	15	0.12	0.02	0.45	0.12	40	5000	40	0.0013
	S12	6.0-0	15	45	25	15	0.1	0.02	0.45	0.1	60	10000	35	0.0011
	S13	6.5-0	10	25	55	10	0.13	0.03	0.48	0.13	100	20000	50	0.0010
	S14	7.0-0	10	25	55	10	0.16	0.03	0.48	0.16	40	20000	45	0.0011
	S15	7.5-0	10	25	55	10	0.13	0.03	0.48	0.13	40	20000	40	0.0012
	S16	8.0-0	10	25	55	10	0.15	0.03	0.48	0.15	30	10000	30	0.0010
	S17	8.5-0	10	25	45	20	0.14	0.02	0.45	0.14	50	50000	50	0.0010
	S18	9.0-0	30	30	18	22	0.13	0.02	0.52	0.13	60	25000	35	0.0012
	S19	9.5-0	30	40	12	18	0.12	0.02	0.55	0.12	55	25000	20	0.0013
	S20	10.0-0	30	42	18	10	0.11	0.01	0.40	0.11	50	5000	15	0.0011
Sediment Property														
	Sediment Type	Diameter	Density	Initial Porosity	Initial Permeability	Compacted Porosity	Compaction	Compacted Permeability	Erodibility					
	Coarse Grained Sand	1.0 mm	2.70 g/cm ³	0.21 m ³ /m ³	500 mD	0.25 m ³ /m ³	5000 KPa	50 mD	0.6					
	Fine Grained Sand	0.1 mm	2.70 g/cm ³	0.3 m ³ /m ³	100 mD	0.15 m ³ /m ³	2500 KPa	5 mD	0.45					
	Silt	0.01 mm	2.65 g/cm ³	0.38 m ³ /m ³	50 mD	0.12 m ³ /m ³	1200 KPa	2 mD	0.3					
	Clay	0.001 mm	2.65 g/cm ³	0.48 m ³ /m ³	5 mD	0.05 m ³ /m ³	500 KPa	0.1 mD	0.15					

		Initial Conditions- GPM Input Parameters												
		Simulation Duration	Sediment Type Proportion (%)				Avg. Water Velocity	Avg. Sediment Velocity	Erodibility	Diffusion Coefficient	Avg. Sea Level	Turbidite Event Interval	Steady Flow Iteration	Sediment Movement
GPM Scenarios (GS)		(Ma- 0a) Years	Sand (Coarse)	Sand (Fine)	Silt	Clay	(m/a)	(m/a)			Interval (m)	(/years)	(/hrs)	Coefficient
	S1	0.02 – 0	25	25	25	25	0.11	0.03	0.35	0.11	30	2500	10	0.001
	S2	0.25 – 0	25	25	25	25	0.15	0.03	0.45	0.15	70	1000	15	0.012
	S3	0.5 – 0	25	25	25	25	0.11	0.02	0.55	0.11	120	1000	20	0.012
	S4	0.7 – 0.05	25	25	25	25	0.08	0.02	0.35	0.08	100	500	25	0.0011
	S5	1.5 – 0	15	35	30	20	0.15	0.04	0.50	0.15	80	5000	20	0.001
	S6	3.0 – 0	50	25	15	10	0.13	0.04	0.50	0.13	70	5000	30	0.0012
	S7	3.5 – 0	50	25	15	10	0.11	0.04	0.50	0.11	70	10000	15	0.001
	S8	4.0 – 0	50	25	15	10	0.13	0.04	0.50	0.13	90	5000	20	0.0015
	S9	4.5 – 0	15	45	25	15	0.1	0.02	0.45	0.1	50	10000	30	0.0012
	S10	5.0 – 0	15	45	25	15	0.12	0.02	0.45	0.12	55	10000	35	0.0013
	S11	5.5 - 0	15	45	25	15	0.12	0.02	0.45	0.12	40	5000	40	0.0013
	S12	6.0 – 0	15	45	25	15	0.1	0.02	0.45	0.1	60	10000	35	0.0011
	S13	6.5 – 0	10	25	55	10	0.13	0.03	0.48	0.13	100	20000	50	0.0010
	S14	7.0 – 0	10	25	55	10	0.16	0.03	0.48	0.16	40	20000	45	0.0011
	S15	7.5 – 0	10	25	55	10	0.13	0.03	0.48	0.13	40	20000	40	0.0012
	S16	8.0 – 0	10	25	55	10	0.15	0.03	0.48	0.15	30	10000	30	0.0010
	S17	8.5 – 0	10	25	45	20	0.14	0.02	0.45	0.14	50	50000	50	0.0010
	S18	9.0 – 0	30	30	18	22	0.13	0.02	0.52	0.13	60	25000	35	0.0012
	S19	9.5 – 0	30	40	12	18	0.12	0.02	0.55	0.12	55	25000	20	0.0013
S20	10.0 - 0	30	42	18	10	0.11	0.01	0.40	0.11	50	5000	15	0.0011	
Sediment Property														
	Sediment Type	Diameter	Density	Initial Porosity	Initial Permeability	Compacted Porosity	Compaction	Compacted Permeability	Erodibility					
	Coarse Grained Sand	1.0 mm	2.70 g/cm ³	0.21 m ³ /m ³	500 mD	0.25 m ³ /m ³	5000 KPa	50 mD	0.6					
	Fine Grained Sand	0.1 mm	2.70 g/cm ³	0.3 m ³ /m ³	100 mD	0.15 m ³ /m ³	2500 KPa	5 mD	0.45					
	Silt	0.01 mm	2.65 g/cm ³	0.38 m ³ /m ³	50 mD	0.12 m ³ /m ³	1200 KPa	2 mD	0.3					
	Clay	0.001 mm	2.65 g/cm ³	0.48 m ³ /m ³	5 mD	0.05 m ³ /m ³	500 KPa	0.1 mD	0.15					

Table 3. Lithofacies classification in the forward stratigraphic model; showing the command used in the property calculator tool in Petrel™.

Lithofacies Classification		
Facies Code	Lithofacies	Command Used in Petrel's Property Calculator
0	Marine Shale	If(Sand_fine>=0.19 And Sand_fine<=0.21 Or Silt>=0.19 And Silt<=0.2 Or Clay>=0.2 And Clay<=0.21 Or Depth_of_deposition>=-82 And Depth_of_deposition<=-78)
1	Muddy Shallow Bay Fill	If(Sand_fine>=0.36 And Sand_fine<=0.38 Or Silt>=0.18 And Silt<=0.2 Or Clay>0.18 And Clay<=0.19 Or Depth_of_deposition>=-30 And Depth_of_deposition<=-20)
2	Sandy Shallow Bay Fill	If(Sand_coarse>=0.65 And Sand_coarse<=0.73 Or Sand_fine>=0.18 And Sand_fine<=0.22 Or Silt>=0.18 And Silt<=0.2 Or Clay>=0.17 And Clay<=0.18 Or Depth_of_deposition>=-3 And Depth_of_deposition<=0)
3	Channel Fill Sandstone	If(Sand_coarse>=0.5 And Sand_coarse<=0.68 Or Sand_fine>=0.23 And Sand_fine<=0.25 Or Silt>=0.17 And Silt<=0.18 Or Depth_of_deposition>=0 And Depth_of_deposition<=2)
4	Lower Shoreface Units	If(Sand_coarse>=0.19 And Sand_coarse<=0.31 Or Sand_fine>=0.19 And Sand_fine<=0.24 Or Silt>=0.4 And Silt<=0.48 Or Clay>=0.19 And Clay<=0.31 Or Depth_of_deposition>=-83 And Depth_of_deposition<=50)
5	Middle Shoreface Units	If(Sand_coarse>=0.32 And Sand_coarse<=0.53 Or Sand_fine>=0.25 And Sand_fine<=0.32 Or Silt>=0.26 And Silt<=0.32 Or Clay>=0.19 And Clay<=0.21 Or Depth_of_deposition>=-38 And Depth_of_deposition<=-12)
6	Upper Shoreface Units	If(Sand_coarse>=0.53 And Sand_coarse<=0.72 Or Sand_fine>=0.28 And Sand_fine<=0.33 Or Silt>=0.16 And Silt<=0.21 Or Depth_of_deposition>=-10 And Depth_of_deposition<=6)
7	Distal Mouth Bar Units	If(Sand_fine>=0.23 And Sand_fine<=0.27 Or Silt>=0.38 And Silt<=0.43 Or Clay>=0.19 And Clay<=0.21 Or Depth_of_deposition>=-95 And Depth_of_deposition<=-80)
8	Proximal Mouth Bar Units	If(Sand_coarse>=0.53 And Sand_coarse<=0.71 Or Sand_fine>=0.27 And Sand_fine<=0.32 Or Silt>=0.16 And Silt<=0.21 Or Clay>=0.06 And Clay<=0.07 Or Depth_of_deposition>=-30 And Depth_of_deposition<=-27)
9	Tide Influenced Sandstones	If(Sand_coarse>=0.53 And Sand_coarse<=0.71 Or Sand_fine>=0.26 And Sand_fine<=0.31 Or Silt>=0.35 And Silt<=0.41 Or Depth_of_deposition>=-5 And Depth_of_deposition<=1)
10	Fluvial Channel Sandstones	If(Sand_coarse>=0.54 And Sand_coarse<=0.56 Or Sand_fine>=0.27 And Sand_fine<=0.29 Or Silt>=0.19 And Silt<=0.21 Or Depth_of_deposition>=-2 And Depth_of_deposition<=2)
11	Coal	Estimated as background attribute
12	Coastal plain fines	If(Silt>=0.31 And Silt<=0.43 Or Clay>=0.31 And Clay<=0.35 Or Depositional_depth>=-100 And Depositional_depth<=-40)
13	Marine Mudstone	If(Sand_fine>=0.36 And Sand_fine<=0.38 Or Silt>=0.4 And Silt<=0.52 Or Clay>=0.45 And Clay<=0.78 Or Depth_of_deposition>=-105 And Depth_of_deposition<=-90)

Lithofacies Classification

Facies Code	Lithofacies	Command Used in Petrel's Property Calculator
0	Marine Shale	If(Sand_fine>=0.19 And Sand_fine<=0.21 Or Silt>=0.19 And Silt<=0.2 Or Clay>=0.2 And Clay<=0.21 Or Depth_of_deposition>=-82 And Depth_of_deposition<=-78)
1	Muddy Shallow Bay Fill	If(Sand_fine>=0.36 And Sand_fine<=0.38 Or Silt>=0.18 And Silt<=0.2 Or Clay>0.18 And Clay<=0.19 Or Depth_of_deposition>=-30 And Depth_of_deposition<=-20)
2	Sandy Shallow Bay Fill	If(Sand_coarse>=0.65 And Sand_coarse<=0.73 Or Sand_fine>=0.18 And Sand_fine<=0.22 Or Silt>=0.18 And Silt<=0.2 Or Clay>=0.17 And Clay<=0.18 Or Depth_of_deposition>=-3 And Depth_of_deposition<=0)
3	Channel Fill Sandstone	If(Sand_coarse>=0.5 And Sand_coarse<=0.68 Or Sand_fine>=0.23 And Sand_fine<=0.25 Or Silt>=0.17 And Silt<=0.18 Or Depth_of_deposition>=0 And Depth_of_deposition<=2)
4	Lower Shoreface Units	If(Sand_coarse>=0.19 And Sand_coarse<=0.31 Or Sand_fine>=0.19 And Sand_fine<=0.24 Or Silt>=0.4 And Silt<=0.48 Or Clay>=0.19 And Clay<=0.31 Or Depth_of_deposition>=-83 And Depth_of_deposition<=50)
5	Middle Shoreface Units	If(Sand_coarse>=0.32 And Sand_coarse<=0.53 Or Sand_fine>=0.25 And Sand_fine<=0.32 Or Silt>=0.26 And Silt<=0.32 Or Clay>=0.19 And Clay<=0.21 Or Depth_of_deposition>=-38 And Depth_of_deposition<=-12)
6	Upper Shoreface Units	If(Sand_coarse>=0.53 And Sand_coarse<=0.72 Or Sand_fine>=0.28 And Sand_fine<=0.33 Or Silt>=0.16 And Silt<=0.21 Or Depth_of_deposition>=-10 And Depth_of_deposition<=6)
7	Distal Mouth Bar Units	If(Sand_fine>=0.23 And Sand_fine<=0.27 Or Silt>=0.38 And Silt<=0.43 Or Clay>=0.19 And Clay<=0.21 Or Depth_of_deposition>=-95 And Depth_of_deposition<=-80)
8	Proximal Mouth Bar Units	If(Sand_coarse>=0.53 And Sand_coarse<=0.71 Or Sand_fine>=0.27 And Sand_fine<=0.32 Or Silt>=0.16 And Silt<=0.21 Or Clay>=0.06 And Clay<=0.07 Or Depth_of_deposition>=-30 And Depth_of_deposition<=-27)
9	Tide Influenced Sandstones	If(Sand_coarse>=0.53 And Sand_coarse<=0.71 Or Sand_fine>=0.26 And Sand_fine<=0.31 Or Silt>=0.35 And Silt<=0.41 Or Depth_of_deposition>=-5 And Depth_of_deposition<=1)
10	Fluvial Channel Sandstones	If(Sand_coarse>=0.54 And Sand_coarse<=0.56 Or Sand_fine>=0.27 And Sand_fine<=0.29 Or Silt>=0.19 And Silt<=0.21 Or Depth_of_deposition>=-2 And Depth_of_deposition<=2)
11	Coal	Estimated as background attribute
12	Coastal plain fines	If(Silt>=0.31 And Silt<=0.43 Or Clay>=0.31 And Clay<=0.35 Or Depositional_depth>=-100 And Depositional_depth<=-40)
13	Marine Mudstone	If(Sand_fine>=0.36 And Sand_fine<=0.38 Or Silt>=0.4 And Silt<=0.52 Or Clay>=0.45 And Clay<=0.78 Or Depth_of_deposition>=-105 And Depth_of_deposition<=-90)

Table 4. Porosity and Permeability estimate in identified lithofacies packages.

Code	Lithofacies	Average NPHI	Density Porosity	Estimated Porosity	KLOGH (mD)
0	Marine Shale	0.17 - 0.45	0.1	0.08 - 0.11	10.02 - 16.1
1	Muddy Shallow Bay Fill	0.17 - 0.42	0.1	0.08 - 0.13	23.85 - 102.3
2	Sandy Shallow Bay Fill	0.07 - 0.52	0.25	0.16 - 0.25	100.0 - 398.7
3	Channel Fill Sandstone	0.04 - 0.15	0.30	0.18 - 0.22	400.01 - 889.7
4	Distal Lower Shoreface	0.04 - 0.19	0.29	0.1 - 0.23	120.5 - 170.3
5	Proximal Shoreface	0.05 - 0.17	0.31	0.17 - 0.24	80.2 - 412.5
6	Upper Shoreface Units	0.09 - 0.11	0.28	0.21 - 0.26	650.2 - 1023.7
7	Distal Mouth Bar Units	0.08 - 0.17	0.27	0.09 - 0.17	170.5 - 223.1
8	Proximal Mouth Bar	0.05 - 0.59	0.12	0.19 - 0.21	130.5 - 314.3
9	Tide Influenced SS	0.14 - 0.46	0.26	0.15 - 0.20	220.0 - 512.6
10	Fluvial Sandstones	0.14 - 0.29	0.21	0.19 - 0.21	180.5 - 691.8
11	Coal	0.24 - 0.53	0.05	0.001	0.001
12	Coastal Plain Fines	0.43 - 0.49	0.06	0.04 - 0.12	5.2 - 34.6
13	Marine Mudstone	0.16 - 0.42	0.1	0.08 - 0.10	6.0 - 15.2

Table 5. Comparison of a) porosity, and b) permeability estimates in original petrophysical model and forward modeling-based porosity and permeability models.

a. Validation Well Position 1							
	Porosity: GPM-Based Model					Porosity: Original Model	
	Depth (m)						
Models	5 m	10 m	15 m	25 m	35 m	Depth (m)	Average Porosity
R14	0.22	0.24	0.16	0.22	0.16	5	0.2
R20	0.16	0.19	0.26	0.18	0.15	10	0.25
R26	0.18	0.17	0.23	0.16	0.19	15	0.27
R36	0.22	0.21	0.19	0.22	0.21	25	0.16
R45	0.25	0.2	0.23	0.22	0.15	35	0.13
R49	0.21	0.17	0.22	0.17	0.18		
Validation Well Position 2							
	Porosity: GPM-Based Model					Porosity: Original Model	
	Depth (m)						
Models	5 m	10 m	15 m	25 m	35 m	Depth (m)	Average Porosity
R14	0.17	0.16	0.24	0.15	0.25	5	0.17
R20	0.21	0.22	0.2	0.21	0.23	10	0.21
R26	0.21	0.2	0.21	0.25	0.24	15	0.21
R36	0.2	0.22	0.21	0.21	0.19	25	0.17
R45	0.22	0.19	0.2	0.19	0.21	35	0.19
R49	0.26	0.24	0.23	0.16	0.21		
b. Validation Well Position 1							
	Permeability_Z (mD): GPM-Based Model					Permeability_Z: Original Model	
	Depth (m)						
Models	5 m	10 m	15 m	25 m	35 m	Depth (m)	Average Perm_Z
R14	163.95	312.38	69.84	310.16	508.2	5	352.74
R20	290.84	315.09	105.66	273.04	200.63	10	312.38
R26	375.92	203.81	166.23	189.92	348.12	15	201.08
R36	418.03	203.27	190.9	168.9	370.56	25	199.76
R45	337.6	412.67	199.66	156.71	305.92	35	508.2
R49	370.89	129.33	291.77	175.53	551.18		
Validation Well Position 2							
	Permeability_Z (mD): GPM-Based Model					Permeability_Z: Original Model	
	Depth (m)						
Models	5 m	10 m	15 m	25 m	35 m	Depth (m)	Average Perm_Z
R14	320.34	336.22	151.08	464.22	132.98	5	6.6
R20	122.66	209.15	161.3	230.58	208.48	10	883.6
R26	151.48	710.07	175.09	384.49	169.48	15	30.3
R36	184.74	344.99	157.08	420.15	136.14	25	496.99
R45	91.44	361.04	77.17	382.85	134.56	35	156.6
R49	134.01	721.73	137.42	636.48	290.06		