

re: “Review of Schneider et al. 2020”, Vincent Verjans, 27 Aug 2020

We appreciate your thorough review and are working to address your concerns. Following below is our response (in blue) to your enumerated major and specific comments, which are italicized here for reference:

1. *The lack of information about the optimisation*

Here are two review criteria of GMD:

- *“Are the methods and assumptions valid and clearly outlined?”*
- *“Is the description sufficiently complete and precise to allow their reproduction by fellow scientists (traceability of results)?”*

I focus first on the optimisation method applied (Section 3.2). The entire optimisation method is described in a single sentence (lines 184-186):

“From our estimated empirical strain rate-versus-depth data, we optimized the previously described densification model coefficients (from A’76, vK’17, and A’10) by applying a regularized least squares algorithm for two stages of densification (above and below $\rho = 550 \text{ kg m}^{-3}$).”

The A’76 model includes 7 different coefficients ($c_1, c_2, c_3, c_4, c_5, \rho_{dm}, \eta_0$), the vK’17 includes 10 ($c_3, c_4, c_5, \rho_{dm}, \eta_0, c_\eta, f_1, f_2, a_\eta, b_\eta$) and the A’10 includes 7 ($c_3, c_4, c_5, \rho_{dm}, k_c^{\rho < 550}, k_c^{\rho > 550}, E_c$). Additionally, in the Results section, the authors mention “adding a constant compaction term” (line 277), which does not appear in any equation and is not explained in the Methods section. Throughout the manuscript, the authors never state which coefficients are subject to the optimisation. Moreover, they mention in Section 4.2 that “we have yet to test in ELM an optimized version of the semi-empirical model”. From my understanding, the “semi-empirical model” is A’10 and they decided to compute the ELM simulations with the original version of A’10 and not the optimised version. In contrast, for some reason, the authors did the ELM simulations with the optimised version (vK’17+) of vK’17. They still provide speculative avenues for a re-parameterisation of A’10 at lines 322-328. These statements are not supported by any quantitative information about a better fit of the optimised A’10 either to observed data or to the strain rates generated from the model of Herron and Langway (1980) (referred to as HL hereafter). Finally, they assert that they optimise A’76 (“we optimized the previously described densification model coefficients (from A’76, vK’17, and A’10)”). However, the only information to be found about the modifications brought to the model is the change of the value of c_5 , but nothing about other parameters and nothing about a performance comparison between the original A’76 and the optimised version.

Clearly, we failed to provide essential details regarding the optimization method used to calibrate firm densification model coefficients. While we experimented with numerous model configurations, the optimization referred to in the manuscript pertains to c_3 , η_0 , and k_c , as referred to above, plus an additional term not mentioned in the manuscript! We appreciate you pointing out this oversight. These coefficients were calibrated using a convoluted least squares algorithm that we must clarify in our revision.

At the time of submission, we did not have ELM simulation results from the optimized version of the A'10 configuration (hereafter A'10+). Ideally, we would test dozens of model configurations in ELM. However the ELM simulations on century timescales are computationally expensive and time consuming. Based on the preliminary results of our statistical modeling, we experimented with a small selection of model configurations (including both vK'17 and vK'17+), which we describe in the manuscript. Our speculative reparameterization of the A'10 model has now been tested in a century-scale ELM simulation, but due to the noted lack of clarity, it is not clear how we arrived at these speculations. Finally, we discovered that the original (A'76) configuration implemented in ELM could be significantly improved by simply modifying c_5 . Therefore, we calculated the optimal value of c_5 using our statistical model, but again, these results are impossible for the reader to fully comprehend due to our lack of clarity.

To address these critical flaws, we are expanding Section 3.2 by adding details regarding the optimization method. We are also adding new results from the A'10+ ELM simulations and are removing the speculative comments toward the end of Section 4.1.

Coming back to the optimisation methodology itself, the authors decide to select annual mean temperatures only below -25°C , but they then proceed to model simulations for grid cells where the annual mean temperature is as high as -20°C (Section 3.3, Figures 2, 3 and 4). It is puzzling that the authors themselves suggest a better approach to selecting mean annual temperatures and accumulation rates, which would be easy to implement (lines 357-361). They also claim to calibrate the models by matching the computed strain rates to the HL strain rates. The issue is that several models use dynamic variables in the strain rate equations: T, σ, P and r_e in Eqs 1, 2, 3, 4, 8, 9 and 10. The HL can provide analytic solutions of strain rates for steady state annual mean temperature and accumulation. The dynamic models require values for the dynamic variables at each time step and for each layer of the firn column. A steady state annual mean temperature does not correspond to all firn layers having the same temperature year-round (temperature still varies seasonally when in steady state). Similarly, accumulation rate still varies seasonally, which means that σ also varies in time for any firn layer (again, even in steady state). Finally, the reader has no information about how r_e is calculated in the computations of these steady state strain rates. The calibration method only considers density profiles for mean annual temperatures less than -25°C so that it will be optimized for dry firn densification. However, ELM simulations provide global results, including warmer grid-cells that are shown in Figs. 2, 3, and 4. In the future, it would be interesting to download output from regional climate models and reanalyses to read into the statistical model. However, such an approach is beyond the scope of this study. Regretfully, we failed to provide enough details about the rest of our statistical model and optimization method including how we calculate the dynamic variables T, σ, P and r_e .

To address these issues, first, we will change the comparisons of ELM results versus Herron and Langway (1980) so that the columns in Figs. 2, 3, and 4 center around mean annual temperatures of $-39, -32,$ and -25°C . This will eliminate the discrepancy in the temperature cutoff referred to above. And second, we are expanding Section 3.2 to include details regarding how dynamic variables (T, σ, P and r_e) are estimated.

Furthermore, the “regularized least square algorithm” is not described. Why not proceed to a simple least square? What is the penalty term? What are the penalised factors? And,

most importantly, which coefficients are subject to the optimisation and what is the range of possible values covered by the optimisation?

Yes, the least squares algorithm is convoluted and totally obscure based on the manuscript alone. None of these reasonable questions are answered and therefore the results cannot be reproduced.

To correct this critical flaw, we are adding to Section 3.2 or an Appendix a more detailed description of the algorithm including all pertinent tuning parameters and coefficients subject to the optimization.

I think that the authors can easily understand that the issues I raise here are concerning with respect to the GMD review criteria.

Yes, the issues you raise show that the originally submitted manuscript lacks the clear explanation needed to make it reproducible. The forthcoming revision will resolve these issues as we describe.

2. *The ELM firn density simulations*

My first concern relates to the data that is used for model evaluation. Why use the cores of Mosley-Thompson et al. (2001) and Lamorey (2003)? And why do the authors average the Greenland cores? They highlight themselves that “variability can be large, particularly across the GrIS” (line 337). Why not compare an observed firn core to the model simulations for the grid cell of the corresponding location? Averaging observed and modelled firn depth-density profiles makes little sense. The authors themselves seem to point out this shortcoming of the study (lines 318-320): “though our analysis with ELM thus far is limited to a generalized comparison with a broad (climate) perspective rather than to a more site-specific comparison against direct observations”. So why was a site-specific comparison not performed?

The evaluation against observations was also challenged by the other reviewers.

To redeem this flaw, we are adding to Section 4 a new analysis that evaluates Greenland Ice Sheet (GrIS) results against the SUMup dataset (Koenig and Montgomery, 2019), as you suggest below. This analysis covers the accumulation areas of the GrIS better than the subset of measurements provided by Mosley-Thompson et al. (2001) and adds a geographical comparison with a quantitative evaluation. By computing a similarity matrix, we categorize measurements nearest in distance and time to corresponding ELM simulations and then assess model accuracy with root mean squared errors (RMSE). This new analysis serves two objectives: one, improve the quality of the model to observational comparison by controlling for space (location) and time (year); and two, offer metrics for quantitatively assessing model accuracy.

Moreover, the data selected are not in line with the objective of the study: improving dry densification schemes. Most of the Greenland grid cells are likely affected by melt, and Siple Dome is an area of Antarctica with relatively high melt rates for the continent. Why don't the authors select data only from dry snow areas (higher accumulation zone of Greenland and more inland regions of Antarctica)? Do the authors know about the extensive SUMup dataset (Koenig and Montgomery, 2019) that includes many more firn cores? The occurrence of melt is clear because there is “formation of ice lenses” (line 241). But no information is provided about the model schemes for simulation of meltwater percolation and refreezing. Moreover, the simulations are performed with atmospheric forcing at very coarse resolution,

which is underlined at lines 342-347 (I mention here that the resolution is not provided in the manuscript). This forces the authors to artificially adjust their evaluation: “this large grid cell remapping lead to a cold bias, resulting in too-slow densification. Therefore, we adjusted our Siple Dome comparisons to include gridcells away from the coast that better represent atmospheric conditions and result in a more realistic density simulation”. Firstly, I would think that grid cells away from the coast should be even colder and thus enhance the cold bias. Secondly, this further underlines the question of why choosing Siple Dome and Greenland firn cores to evaluate the models. This choice brings in problems related to the adequacy of the model forcing, which makes it very difficult to disentangle firn model deficiencies from errors due to inadequate forcing.

To restrict our evaluation to grid-cells without melt, we originally removed ELM columns that experience mean annual temperatures of greater than -25°C . Regarding the Siple Dome case study, we were mistaken about the “...cold bias, resulting in too-slow densification” (line 345). Rather, ELM has a “warm” bias there, resulting in too much densification, so we adjusted the ELM comparison inland, which is more representative of the actual Siple Dome climate and resulted in better agreement with the observed density profiles. Your point about the difficulty in evaluating coarse resolution ELM grid cells to point based measurements is why we generally fall back to the model of Herron and Langway (1980) to evaluate model performance. The HL empirical model can be applied to gridcells that experience faulty atmospheric input data. The comparisons plotted in Figs. 2, 3, and 4 control for mean annual temperature and accumulation to help disentangle firn model deficiencies from inaccurate atmospheric forcing. As mentioned above, we now have a much more comprehensive evaluation against the SUMup (measurement) dataset (Koenig and Montgomery, 2019).

Because both original evaluation methods are questionable, we are adding to Section 4 another evaluation that better controls for geographic location and time and also provides RMSEs with reference to SUMup observations (Koenig and Montgomery, 2019). This new evaluation indicates that our optimization method improves upon the snowpack and firn compaction model currently in CLM for most model grid-cells across the GrIS.

The spin-up period is taken to be 260 years. This is likely too short for low-accumulation grid cells (thus most of the dry snow zone) to build a full firn layer (i.e. until ice density is reached). The authors should thus support their statement (lines 226-227) that the profiles “averaged from the final 100 years of simulation results” (thus starting the averaging only after 160 years) are “steady-state density profiles”. For example, after 160 years of an accumulation rate of $0.07\text{ m w.e. yr}^{-1}$ (the limit assumed for warm dry snow zones in Section 3.2), only 11 m w.e. have accumulated, which corresponds roughly to a firn column of 20 m. I doubt that this represents a steady state. As shown in Figure 4 ($T=-34^{\circ}\text{C}$), firn that is 100 years old is only at 600 kg m^{-3} density, showing that the firn layer is most likely not in steady state after 160 years and not even after 260 years. Concerning the fresh snow density, the A’76 model calculates surface densities by itself, while vK’17 and vK’17+ use a fresh snow parameterisation (not given in the manuscript...). But how is fresh snow density calculated for the A’10 model? The prognostic equation for r_e should also be given or referred to.

We are aware of the difficulties pertaining to fully spinning up the firn column in low accumulation grid-cells that represent most of the Antarctic Ice Sheet (AIS). That is why we only evaluate firn columns that exceed a thickness of 60 m. By filtering out columns that have not reached this threshold, we are only evaluating relatively high accumulation grid-cells mostly

representing the GrIS. This is also why we only show relatively warm grid-cells in Figs. 2, 3, and 4, i.e., because most of the grid-cells that have relatively low accumulation rates are much colder and have not reached a depth of 60 m.

To address concerns raised above, we are adding a demonstration showing the time evolution of the simulated firn density profiles used for evaluation. Because the deep density trends over the last 100 years are statistically indistinguishable from zero, with variations occurring almost exclusively in the top 10 m, our 260 year spin up is adequate for evaluating grid-cells that accumulate at least 0.2 m SWE yr^{-1} , covering most of the GrIS and parts of the AIS.

We will also add to the revision that the “A’76” model configuration uses a constant new snow density of 50 kg m^{-3} . And we will add the new snow density parameterization by van Kampenhout et al. (2017) to the revised Section 3, where we provide a streamlined introduction of each of the ELM experiments evaluated in Section 4. The prognostic equation for r_e is included in the Snow, Ice, and Aerosol Radiative Model (SNICAR) referenced in lines 75-77. We will add r_e after “ice effective grain size (from SNICAR)” for clarity.

The approximation of vertical strain rates (line 244) is also unclear. This raises the same questions that I mentioned above about the dynamic variables when assuming a steady state. In my view, the authors should include a detailed explanation about how the steady state strain rates of A’76, A’10, vK’17 and vK’17+ are calculated. This holds for both the calibration step as for the ELM simulations (i.e. the values appearing in Figure 4).

We agree that it can be difficult to follow the nuances associated with all the different model configurations, including both within the statistical model and the ELM experiments. However, all of strain rates calculated in our study are from the equations provided in Section 2.

As we attempt to mitigate this confusion, we are following your suggestion to expand the description of the statistical model. We hope our revision of Section 3 will reduce obscurity and, in addition to adding details discussed in previous responses, provide readers with clearer steps on how to reproduce our statistical modeling results and approximated strain rates from ELM simulations.

When analysing and evaluating the results, there is a severe lack of quantification. This holds for both the Equilibrium climate simulations and the Twentieth century climate simulations. I give some examples:

- *“the semi-empirical model improves the density profile” (line 260): improves with respect to what (I guess that the authors mean A’10 improves the density profile with respect to A’76)? And the statement of “improvement” should not be based on a mere visual comparison of Figures 2 and 3. Moreover, it should be clarified that the authors evaluate the model results against results from HL, which is not a guarantee of model accuracy.*

You guess correctly, though you should not have had to guess at all.

We quantify improvements with our new analysis, which computes RMSEs of ELM density profiles with reference to the nearest measurements from the SUMup dataset (Koenig and Montgomery, 2019). This new analysis supports our previous qualitative claims, which will be moved into a new “Discussion” section (5). We hope that our new format, having two separate “Results” and “Discussion” sections, will allow readers to better interpret our quantitative results without confounding influences from our more

speculative remarks.

- “Densification tapers-off at lower densities (around 450 kg m^{-3}) for colder climates, a temperature-dependent effect enhanced with the model from Arthern et al. (2010).” (lines 248-250): from Figures 3 and 4, it is not obvious that this effect is [s]tronger in A’10 than in vK’17+. The enhancement of the effect should thus be quantified.

Agreed.

We will quantify this statement, move it into the discussion, and further qualify it by replacing “...effect enhanced...” with “...effect slightly enhanced...”

- “A lower model variance occurs when it does not covary with the empirical model. This effectively reduces a model’s prediction risk if it does not also result in an increased bias.” (lines 303-304): If the authors discuss about the bias of models, they can simply add a “Bias” column in Table 2.
- “both models show improvement compared to their original counterparts (ELM v1 and CLM).” (lines 337-338): this is impossible to evaluate for the reader because (1) only the results of A’10 and vK’17+ are shown in Figure 6, and not the ones of their so-called “original counterparts” (which are A’76 and vK’17 I suppose), and (2) there is no quantitative evaluation of the models’ performance with respect to the observed data (e.g. RMSE, bias, etc.).

Our lack of a sufficient quantitative analysis is the fundamental motif in the open discussion of this study.

Our revision includes an expanded Section 4 that streamlines the presentation of ELM experimental results and also includes RMSEs computed with reference to the SUMup measurement dataset (Koenig and Montgomery, 2019).

- “Encouragingly, our simulation results compare well with firn density measurements and indicate an improved capability in the ELM.” (lines 348-349): same remarks as for the previous point.

Again (not surprisingly), our lack of a sufficient quantitative analysis is the fundamental motif in the open discussion of this study.

In our revision we expand Section 4, which now streamlines the presentation of ELM experimental results and also includes RMSEs computed with reference to the SUMup measurement dataset (Koenig and Montgomery, 2019).

3. The novelty and objective of the study

GMD review criterion:

- Does the paper present novel concepts, ideas, tools, or data?

Firn model optimisation has been addressed in numerous studies over the recent years (e.g. Ligtenberg et al., 2011; Kuipers Munneke et al., 2015; van Kampenhout et al., 2017; Smith et al., 2020; Verjans et al., 2020). Four of the studies mentioned have already investigated the optimisation of the model of Arthern et al. (2010), for Antarctica, Greenland or both. An easy and straightforward way to improve the ELM would be to implement the parameterisations developed in these studies. If the authors want to address the same problem, they should propose a new, original method. However, in contrast to the existing literature, they do not calibrate the model of Arthern et al. (2010) with observations but with HL-computed strain

rates. And, as mentioned above, it is unclear to me how they calibrate a dynamic model to steady state strain rates. They should justify why their methodology is better suited to their objective than using what other researchers have already accomplished. Moreover, the objective stated in the conclusion of improving the capability of the ELM to better simulate refreezing rates in firn is not in line with the study itself. The focus of the calibration is on dry firn densification and does not support the statements at lines (377-381): “With an evaluation of the simulation of dry firn densification, we have optimized the ELM firn model for future studies of the impacts of liquid water on firn density and SMB. Ultimately, this study seeks to enable better predictions of SLR as a direct result of surface melt and mass loss from the GrIS.”

Currently, as far as we know, there is only one other ESM that attempts to include snowpack processes at this level of detail (CESM), which we discuss and compare to in our study. We are not aware of any other firn parameter optimization studies that provide a complete set of values that can simply plug into ELM. Furthermore, because ELM includes two dry snow densification processes (destructive metamorphism and overburden pressure compaction), assimilating published firn densification models into ELM requires calibrating their coefficients to account for both compaction terms. What is also novel about our study is that we experiment with combinations of existing snowpack and firn densification parameterizations. We optimize using the model of HL’80 as calibration data because it can more easily be applied to regions where observations are limited. Even with a complete observational data product, density profiles would still need a smoothing filter to improve the numerical stability of the vertical differentials used in the calculation of strain rates. We use this method to optimize strain rate data to steady strain rates, though we failed to provide adequate documentation regarding the pressure, temperature, and grain size inputs to the model.

We aim to address this problem by providing more details regarding the statistical firn model used to calculate pressure and temperature profiles. We are also removing the misleading quote stated above. Our revised manuscript will also include some additional discussion in the introduction to stress that the goal here is not to make the best firn model ever and stick that in our ESM. Rather, we need a model that balances physical realism against computational cost.

If the authors do want to better capture liquid water effects, they should focus on this very challenging topic by studying the mechanisms of wet firn compaction, meltwater percolation and refreezing.

Wet firn densification is outside the scope of this study. It could be the topic of a future study, but it would be premature to try to improve the wet processes until the dry processes have been fully vetted.

4. *The clarity of the manuscript*

GMD review criterion:

- *Is the overall presentation well structured and clear?*

It is very difficult for the reader to understand the different steps of the study. The authors alternate between different ways to refer to a same thing. For example, the A’10 model is sometimes referred to as “A’10” and sometimes as “the semi-empirical model”, the vK’17 model is sometimes referred to as “vK’17”, sometimes as “the CLM” and sometimes as the “Snowpack model” (see Figure 5). Similarly, it is never clearly stated that the “empirical

strain rates” correspond to the ones computed with the HL. A first, simple way to improve the clarity would be to consistently use the terms A’76, A’10, vK’17 and vK’17+ throughout the manuscript, including in the captions of the Figures. In line with this, the Section 3.1 should be split in four subsections that clearly detail each of these four models instead of subsections presenting equations which are subsequently assigned to the models in a confusing way.

We appreciate these constructive comments.

As a first step, we are following your suggestion regarding Section 3.1. Second, we will remain consistent in our labeling of ELM experiments throughout the manuscript, which entails replacing the implicit descriptors you mention above.

The Figures and Tables also lack clarity. In Figures 2 and 3, why are high values of accumulation only shown at depths greater than 15 m? Even in a high-accumulation climate, there will always be a shallow and a deep part of the firn column. And how were the surface density values chosen for the HL-computed profiles? The caption of Figure 1 mentions that the new firn model “can extend as deep as 80 m”, whereas it is always presented as a “semi-infinite” grid in the text. Which of these two statements is true? In Figure 4, why are there points without a vertical error bar? And why do some points have a horizontal error bar (age should be well-determined for any firn layer of any model run)? In Table 1, equation numbers could be provided for each model to know which equations apply to which model. In Table 2, the model names should be used in the column “Densification model” instead of the mechanisms applied and the variable for which the statistics are calculated should be specified in the caption (presumably strain rate values). Improving the structure of the manuscript could possibly help decrease the degree of confusion for the reader when trying to understand the study.

The empirical model from Herron and Langway (1980) gives density profiles that do not vary with accumulation rate for depths less (more shallow) than the critical depth, i.e., where density is less than 550 kg m^{-3} . In Figs. 2 and 3, the high values of accumulation are obscured by the lowest value. The surface densities range from 300 to 380 kg m^{-3} , which are appropriate for ice sheets (van Kampenhout et al., 2017; Herron and Langway, 1980).

The maximum depth of the new snowpack model is a recurrent point of confusion for readers. While the new vertical grid has a bottom most layer without a bound on its maximum thickness, we choose 60 m as the limit for most of our analysis because that is roughly how far below the surface the 15 layers of finite thickness will extend when they reach their maximum thicknesses (Figure 1). Therefore, the grid is truly semi-infinite, and we choose to exclude the semi-infinite layer from (and thus restrict the vertical extent of) our analysis because the model cannot resolve dynamic variables deeper than 60 m at a vertical scale appropriate for simulating densification (see also response to anonymous referee 2).

In Fig. 4, it appears that some of the vertical error bars are obscured. We speculate this is caused by poor rendering of vector graphics. Horizontal error bars represent grid-cell standard deviations grouped by their Eulerian reference layer (i.e. fixed depth). A layer z_i , e.g., across the global land surface domain represents firn that varies significantly in age. This results in a blurred analysis that we account for with horizontal error bars.

To improve clarity, we changed the wording in the caption of Fig. 1, which replaces “80 m” with “semi-infinite.” Next, we will try using a different image format (or change the line styles) for Fig. 4 to elucidate the vertical error bars. The revision will more clearly explain the meanings of all the symbols (axes, limits, etc...) in the caption. Finally we appreciate

your suggestions on how to clarify information given in Tables 1 and 2 and will follow them accordingly.

- *line 2:*
Change “consist” to consists.
Corrected.
- *line 15:*
I doubt that any paleoclimate study uses Earth System Models to determine pore close off depth and timing.
They are not able to using current Earth System Models, further motivating this study.
- *line 25:*
Repetition of “coupled”.
We will rephrase this sentence in our revision.
- *lines 32-35:*
As far as I understand, there is a contradiction between “does not yet exist” and explaining the implementation of the advanced firn model in the CLM.
The advanced firn model in CLM (version 5) limits its snow water equivalent (SWE) depth to 10 m, which likely corresponds to a maximum allotted firn thickness of less than 30 m. Furthermore, the CLM firn densification model is not valid for densities greater than 550 kg m⁻³ (stage 2).
We will elaborate this point in our revised “Discussion” section (5).
- *lines 47-48:*
Change “those predicted by Herron and Langway (1980)” to “those predicted by the model of Herron and Langway (1980).” Moreover, I suggest using a consistent way to refer to this model (e.g. HL’80).
Thank you; we will follow these suggestions.
- *lines 52-55:*
Add an explanatory sentence about the fact that snowpack models and firn models also have a different vertical scale of application.
Nice suggestion. We will do so accordingly.
- *Section 2.1:*
Provide units of all the variables and quantities presented. This will make clear that there are some unit inconsistencies in some of the equations (which I give below).
We will check all equations and variables for consistency and add units accordingly.
- *Equations 1, 2, 3 and 4:*
The variables $\dot{\epsilon}$ and $\left(\frac{1}{\Delta z} \frac{\partial \Delta z}{\partial t}\right)$ are equivalent to each other as far as I understand. Use either one of the two notations.
We will add a statement of their equivalence and refer to their following quantities as $\dot{\epsilon}$.
- *line 78:*
Specify if $\left(\frac{1}{\Delta z} \frac{\partial \Delta z}{\partial t}\right)_{dm}$ is also considered in CLM (v5).
Thank you for pointing this out. We will clarify that the destructive metamorphism (dm) parameterization is also used in CLM (v5).

- *line 82:*
 In CLM(v5), $c_\eta = 358 \text{ kg m}^{-3}$ and f_2 was set to 4. Only f_1 accounts for the effects of liquid water and not $c_\eta/(f_1 f_2)$.
 In CLM(v5), $c_\eta = 450 \text{ kg m}^{-3}$
<https://github.com/ESCOMP/CTSM/blob/master/src/biogeophys/SnowHydrologyMod.F90>, line 3766). This conflicts with van Kampenhout et al. (2017), who specify $c_\eta = 358 \text{ kg m}^{-3}$. Because we adopt the same constant value for f_2 , the quotient c_η/f_2 can be lumped together as a constant coefficient, and the entire expression $c_\eta/(f_1 f_2)$ still depends on the liquid water content via f_1 .
- *line 89:*
 The characteristic depth is not “a single valued proxy for a given site’s full density profile” but only for the upper density profile.
 We will replace “full” with “upper.”
- *line 89:*
 No s at “stages.”
 Thank you. We will correct.
- *line 89:*
 Add here the explanatory sentence about why models assume a two-stage densification process.
 Done.
- *line 92:*
 The sentence “Empirical firn densification models typically employ analytic functions that assume a steady-state density profile” is not true. Only the model of Herron and Langway (1980) and a few others provide analytic functions but almost all of the recent firn models are dynamic models.
 It is our understanding that the older “empirical” models use statistical methods (linear regression) to calibrate analytical formulations directly to density measurements.
 We will replace the above sentence with “Empirical firn densification models have historically employed analytic functions that assume a steady-state density profile.”
- *line 93:*
 Rephrase.
 Will consider.
- *Equation 5:*
 This equation is erroneous. The units of the left- and right-hand sides do not match. The correction is: $w(z) = \frac{A}{\rho(z)/\rho_w}$, where ρ_w is the density of water (1000 kg m^{-3}).
 Here is our dimensional analysis:
 $w(z) [\text{m s}^{-1}] = A/\rho(z) [\text{kg m}^{-2} \text{ s}^{-1} \times \text{kg}^{-1} \text{ m}^3] = [\text{m s}^{-1}]$, where A is the accumulation rate in terms of $\text{kg m}^{-2} \text{ s}^{-1}$ equivalent to mm of snow water equivalent (SWE) per second.
 We will correct line 100, replacing “(... SWE per year)” with “(... mm SWE s⁻¹).”
- *Equation 7:*
 Again, this equation is erroneous and there is a unit inconsistency. The correction suggested above, fixes the error.

Our correction above fixes the error.

- *Equation 8:*

The variable P is defined here as the “overburden pressure”, which makes it equivalent to σ . I suspect that this variable corresponds to the P as defined in Equation 9, which should be called the “grain-load stress.” I underline here that in the model of Arthern et al. (2010), it is really σ that is used and not the grain-load stress. The authors should explain why they differ from the original model of Arthern et al. (2010) on this point.

These are some interesting points that are confusing in the literature. Here, we reserve σ for the vertically-integrated (column) areal density [kg m^{-2}] and P for the overburden stress [Pa]. Because of inconsistent use of tuning coefficients in the literature, this subtle distinction is only apparent in dimensional analysis. In CLM(v5), for example, van Kampenhout et al. (2017) are indeed referring to σ , which requires being multiplied by the acceleration due to gravity (g) [m s^{-2}] to convert this quantity to “overburden pressure” as a stress. These conversions get washed out by tuning coefficients, which are (again) inconsistent across the literature.

In line 158, we do refer to the “grain-load stress” as you suggest. You are correct, however, in that Arthern et al. (2010) are referring to the “overburden pressure” in their model and not the “grain-load stress” as we adopt in our implementation of their model. Our choice is motivated in the discussion of Arthern et al. (2010), where they state that their model results in densification that occurs too slowly over the interior ice sheets. Therefore, we apply the “grain-load stress” instead of “overburden pressure,” as in eq. (9), because that increases the magnitude of their modeled strain rates by approximately a factor of 2. Furthermore, after applying our offline calibration method, the distinction between these two approaches (neither of which represents a mechanically sound method, by the way), is of second order importance.

For clarity, we will add a brief discussion (similar to above) after eq. (9) motivating our choice of “grain-load stress” P .

- *lines 176-177:*

Specify that the “plausible firn density-versus-depth profiles” were computed with the different models and the HL.

We already mention (albeit compactly) that the “plausible firn density-versus-depth profiles” were computed using the model of Herron and Langway (1980) in lines 182-183.

- *Section 3.2:*

Why do the authors decide to draw annual temperatures from a distribution representative of the global Earth climates instead of the polar climates? Is the objective to have much more values close to $T = -25^\circ\text{C}$? This should be clarified.

Because we will soon study what happens when dry firn starts to melt, we optimize our calibration routine to be more representative of values close to $T = -25^\circ\text{C}$. In the future, as discussed above, we could drive the statistical model with input data from a regional climate model or polar reanalysis, however, this is beyond the scope of our present study.

In our revision, we will specify our motivation for favoring values close to $T = -25^\circ\text{C}$.

- *How are all these values decided:*

We reviewed the work by Herron and Langway (1980) and arrived at the following values

based on the domain of observations applied in their study.

- -25°C as a threshold (whereas ELM simulations involve warmer sites)

This is a rough estimate partitioning Greenland's dry snow zones that we obtained from Cuffey and Paterson (2010).

We will add the reference.

- -51°C as limit between low- and high-accumulation sites (many sites can have $T > -51^{\circ}\text{C}$ and $A < 0.07 \text{ m SWE yr}^{-1}$)

Because we did not want unrealistically high accumulation values for $T < -51^{\circ}\text{C}$, we selected this limit based on observations tabulated by Herron and Langway (1980). Our statistical model does not consider low accumulation values for $T > -51^{\circ}\text{C}$ because they are not well represented by the model of Herron and Langway (1980).

- surface density values between 300 and 380 kg m^{-3}

These limits were deduced from the range of values selected by Herron and Langway (1980) most applicable to our domain of interest (i.e., dry snow zones).

- line 196/

Specify the resolution of “coarse-resolution”.

The horizontal resolution of the ne11 grid is 2.8° .

As requested, we will provide the nominal resolution when referring to ne11.

- line 197:

What does “(an “I-compset” at ne11 resolution)” mean?

This is our Earth system modeling jargon, meaning stand-alone land surface model (i.e. uncoupled ELM) at 2.8° horizontal resolution.

We will add the ne11 nominal resolution.

- line 199:

Change “January 1st” to “January 1st 1901.”

Okay.

- line 206:

Define “restart runs”. In general, it is good practice to define any term used that may not be straightforward to everyone reading the study.

“Restart runs” are in parentheses. They are defined in the corresponding sentence.

- Table 1: *Add a column for ρ_{dm} values. Add another column that indicates which equations apply to which model, with the corresponding equation numbers (see Major Comment 4).*

Thank you for the helpful suggestion. We will update Table 1 accordingly.

- lines 209-215:

All the details provided about the observational data should be given in a separate subsection. Are the sites of measurements affected by surface melt? And are these firn core measurements open access?

We will move this paragraph into its own subsection and offer our perspectives on the applicability and quality of these publicly available measurements.

- Section 4:

There are a lot of speculative statements in this section. I suggest splitting it into a section

Results and a section Discussion, so that the reader can distinguish between model results and the thoughts of the authors.

Thank you for this excellent suggestion. We appreciate your comment as it is clear you have put some thought into how the manuscript can be better structured to benefit readers.

We are adding a separate “Discussion” section (5). This entails moving most of the material currently in Section 4 into the new Section 5 and adding an updated analysis to the beginning of Section 4. We think you will appreciate seeing your suggestions realized in our revision, which will resolve many of the issues you bring to our attention.

- *line 219:*

Change “accumulations” to “accumulation rates.”

Done.

- *lines 217-222:*

Here, the entire methodology is again defined. This can be confusing for the readers. For example, I suggest rephrasing the sentence “To improve the accuracy of our firm model simulations, we optimize compaction terms against empirical strain rates using statistical modeling” because this was the point of a previous section.

We disagree that the “entire methodology is again defined.” Rather, this is a very brief summary of the methodology and introduces our analysis below.

This paragraph will be trimmed and moved in our revision.

- *line 229:*

Does the statement “the mean annual temperature is within a couple degrees of -25 °C” refer to the results of Figure 2 for $T = -27^{\circ}\text{C}$? If so, it would be clearer to give the exact mean annual temperature value.

We will give the exact values of the mean annual temperatures instead of the imprecise statement above.

- *lines 236-237:*

“These simulations demonstrate a stronger effect of temperature on densification rates, resulting in more variation in density with depth (Fig 3).” I think that this sentence means more variable depth-density profiles according to the mean annual temperature. If so, it should be rephrased.

This sentence is being moved into and rephrased in our new “Discussion” section.

- *lines 242-243:*

Specify which models use the “better fresh snow density parameterization”.

We will add discussion of which models result in better fresh snow density.

- *lines 260-261:*

Strange use of commas.

We removed the comma after “ELM”.

- *line 264:*

What is “over-densification”?

Unrealistically high densification rates.

Here, and throughout the manuscript as necessary, we will refrain from using this term in favor of a better description.

- *line 265:*
The notion of “density profiles that vary too weakly with depth” should be replaced with the one of density that increases too weakly in depth.
We will make the relevant changes.
- *line 265:*
What does “Their” refer to?
Arthern et al. (2010). We will specify.
- *lines 274-277:*
Are these the only coefficients included in the optimisation? Or the only ones that were decided to be changed? See Major Comment 1.
Yes, these are the only coefficients included in the optimization, as we describe in our response to your first major comment.
We will clearly state these details in our expanded statistical modeling subsection.
- *line 276:*
The coefficient f_2 is related to the grain radius (see Vionnet et al., 2012). If the ELM calculates grain size, f_2 can be calculated accordingly. It is crucial that the authors clarify why they decide not to follow the original formulation of f_2 (as a function of grain size) but to consider it as a pure tuning factor. This is all the more relevant since it is emphasised throughout the manuscript that firn models need to account for microphysical features.
Here, we follow van Kampenhout et al. (2017). But perhaps we should experiment with the original formulation from Vionnet et al. (2012) in the future.
For now, we will add a discussion regarding this parameter, including its original value, the value adopted by CLM(v5), and in our study as a pure tuning factor.
- *line 286:*
Specify that the density model coefficients are calibrated to the HL-computed strain rates.
We will specify this detail.
- *lines 288-290:*
Specify the variable for which the statistics are calculated. See Major Comment 4.
This sentence refers to the strain rate $\dot{\epsilon}$, which we will specify.
- *lines 293-298:*
How do the authors explain these results?
These results suggest that the overburden pressure dependence, which is part of most firn densification models, complicates its calculation of realistic strain rates. A simpler, temperature-density dependent equation – while originally used for calculating compaction due to destructive metamorphism – can be modified so that it results in densification rates more highly correlated with those predicted by the model of Herron and Langway (1980). This is a surprising result that challenges the use overburden pressure in firn densification models that depend mostly on their functions of density.
- *Table 2:*
The mention to eq. (5) is an error because it is not the one for destructive metamorphism and c_5 does not appear in this equation.
Thank you for bringing this to our attention.

We will replace “eq. (5)” with “eq. (2)”, which is the correct equation.

- line 300:

The value of $R^2 = 0.67$ is valid only for strain rates in the second stage. Note also that Table 2 shows $R^2 = 0.66$.

We will add that $R^2 = 0.50$ for the first stage.

A difference in R^2 of 0.01 is negligible and a result of repeating experiments that generate input values at random.

- line 303:

“A lower model variance occurs when it does not covary with the empirical model”. This statement sounds like a general statement, but I believe that it is applicable only to the results of this specific study.

We will rephrase as “These smaller variances indicate that our experimental model does not covary with values derived from the model of Herron and Langway (1980)”

- line 306:

What does “these results” refer to? The paragraph above is about variances in the compaction rates.

Will change “these results indicate” to “our covariance analysis shows”

- lines 306-307:

“negative correlations between overburden pressure and empirical strain rates”: this is explained by higher overburden being applied to deeper firm, which is at higher density and thus compacts less. A simple explanation could be provided to the reader.

- lines 311-312:

Note that Equations 3, 4 and 8 are directly dependent on density.

- lines 320-328:

Are these results or speculations? Where do these values come from? And did the “statistical computing” focus only on these specific coefficients of A’10? Such conclusions should not be stated without quantitative results to support them. I emphasise again the need to clarify the optimisation method and its results.

These are results from our statistical computing, which we speculate will improve the A’10 experimental results. We will include results from this ELM experiment (“A’10+”) in our revision, thus reducing the speculative nature of these results.

We are reorganizing and expanding Section 3 so that our statistical model is more transparent, expanding Section 4 so our ELM experimental results are more streamlined, and adding a “Discussion” section (5) to signal a clear distinction of results from our more speculative remarks regarding future model development.

- lines 331-332:

Why did the authors decide to use the optimised vK’17 but not the optimised model of A’10?

We had not completed ELM experiments using the optimized model of A’10 (“A’10+”).

With the A’10+ ELM experiment now complete, we will provide these results in our revised Section 4.

- line 333:

Typo “the the”

Corrected.

- *line 341:*
“ne11” is not defined.
The horizontal resolution of the ne11 grid is 2.8°.
As requested, we will provide the nominal resolution when referring to ne11.
- *line 350:*
The statement “we should focus on the near surface layers, as they contain the primary SMB components” should be clarified.
We removed this sentence.
- *line 351:*
“it could be necessary to model the upper most 20 m”: Is this figure of 20 m supported by the results? If not, references should be provided.
We removed this sentence.
- *lines 352-353:*
“the optimized version (vK’17+) is likely the better choice for implementation into the next major release of the E3SM”: Again, it is unclear if this is speculative or if the results provide strong evidence for this. And how would an optimised version of A’10 compare to vK’17+?
With ELM simulations of an optimized version of A’10 now complete, we can offer a more thorough discussion of the optimized model comparisons. In our revision, we include results from both of these particular ELM experiments in the expanded Section 4 and their respective RMSEs calculated with reference to the SUMup dataset (Koenig and Montgomery, 2019). Furthermore, the comparison of A’10+ and vK’17+ model accuracy is the focus of a paragraph in our revised “Discussion” section.
- *lines 353-355:*
A low bias is usually preferable to a high bias. What is meant by this sentence?
We removed this poorly worded sentence.
- *line 364:*
“In the near future, this could be the entire GrIS”: What is the “near future”? And references should be provided.
We will replace “In the near future” with “By 2100” and cite Machguth et al. (2016) as recent evidence suggesting that feedback processes will likely cause the GrIS melt extent to expand rapidly in a warming climate.
- *Conclusions:*
See Major Comment 3.
Regarding lines 377-381, we replace “... have optimized.... Ultimately, this study seeks to enable better predictions of SLR as a direct result of surface melt and mass loss from the GrIS.” with “...prepared.... This study marks progress toward better predictions of SLR caused by surface melting of the GrIS.”
- *line 370:*
Change “with steady-state empirical models” to “with a steady-state empirical model.”
Thank you; we will change the text accordingly.
- *lines 370-371:*

I do not believe that the analysis is “similar to that by van Kampenhout et al. (2017) for CLM.”

We will replace “analysis” with “model development”, which parallels the work of van Kampenhout et al. (2017) in CLM.

- *Code and data availability:*

The availability of the firn core measurements should be provided in this section.

We will seek guidance from the topical editor regarding the policy associated with previously published data.

- *Figures:*

See Major Comment 4. Also, in all Figure captions and legends, the models should be consistently referred to as A’76, A’10, vK’17 and vK’17+.

Please see our response to Major Comment 4.

- *Figure 1:*

The vertical extension of “80 m deep” (caption) contradicts the “semi-infinite” stated in the text.

We changed the wording in the caption to provide clarity on the semi-infinite bottom layer and removed the mention of 80 m.

- *Figures 2 and 3:*

The vertical/horizontal lines at $\rho = 550 \text{ kg m}^{-3}$ and $\rho = 550 \text{ kg m}^{-3}$ can be removed to improve the clarity of the Figures.

Why do high accumulation rate values appear only at great depth?

See our response to Major Comment 4.

- *Figure 4:*

“for various plausible accumulation rates (0.11–0.50 m SWE yr⁻¹)”: Why is 0.11 m SWE yr⁻¹ chosen as lower limit?

This limit was chosen as a suitable value based on observations referred to by Herron and Langway (1980).

The construction of the horizontal error bars is not clear to me.

We described the horizontal error bar above, but will clarify this in a revised figure.

- *Figure 5:*

Do these plots show the results of A’10 (top) and vK’17 (bottom)?

“A’10” and “vK’17” refer to implementations in ELM. Fig. 5 shows results from the combined destructive metamorphism plus overburden pressure compaction equations after calibration.

The revised structure will clarify the distinction between offline firn model configurations targeted in the optimization versus similar implementations in ELM experiments.

For a direct look at how we address these comments, please see the forthcoming revised manuscript.

Sincerely,

A handwritten signature in black ink that reads "Adam Schneider". The signature is written in a cursive style with a large initial 'A' and a long, sweeping underline.

Adam M. Schneider et al.