Validation of the PALM model system 6.0 in a real urban environment; case study of on Prague-Dejvice, Czech Republic

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Abstract. In recent years, The PALM 6.0 model modelling system has been rapidly developed in the recent years with respect to developing its capability to simulate physical processes within urban environments. In this regard, it includes e.g. Some examples in this regard are energy-balance solvers for building and land surfaces, a radiative transfer model to account for multiple reflections and shading, as well as a plant-canopy model to consider the effects of plants on the flow (thermo)dynamicsof

- 5 the flow, and a chemistry transport model to enable simulation of air quality. This study provides a thorough evaluation of modelled meteorological, air chemistry, and ground and wall-surface quantities against dedicated in-situ measurements taken in an urban environment in Prague, Dejvice, Czech Republic. Measurements included e.g. monitoring of air quality and meteorology in street canyons, surface temperature scanning with infrared eamera-cameras, and monitoring of wall heat fluxes. Large-eddy simulations (LES) using the PALM model driven by boundary conditions obtained from a mesoscale model were performed
- 10 for multiple days within two summer and three winter episodes that are characterized by different atmospheric conditionswere performed with the PALM model driven by boundary conditions obtained from a mesoscale model.

For the simulated episodes, the resulting temperature, wind speedand concentrations of chemical compounds, and chemical compound concentrations within street canyons agreed well with the observations, except show realistic representation of the observed state, except that the LES did not adequately capture nighttime cooling near the surface at for certain meteorological

- 15 conditions. In some situations , less insufficient turbulent mixing was modelled, resulting in higher near-surface concentrations. At most of the surface evaluation points the simulated wall-surface temperature agreed fairly well with the observed one regarding its absolute value as well as daily amplitude surface temperature reproduces the observed surface temperature reasonably well for both absolute and daily amplitude values. However, especially for the winter episodes and for modern buildings with multi-layer walls, the heat transfer through the wall walls is partly not well captured, leading to discrepan-
- 20 cies between the modelled and observed wall-surface temperature. Furthermore, we show that model results depend the study corroborates model dependency on the accuracy of the input data, particularly. In particular, the temperatures of surfaces affected by nearby trees strongly depend on the spatial distribution of the leaf area density, land-surface temperatures at grass surfaces strongly depend on the initial soil moisture, or wall-surface temperatures depend on the correct prescription setting

of wall material parameters, though these parameters are often not available with and concentrations on detailed information

- 25 on spatial distribution of emissions, all of which are often unavailable at sufficient accuracy. Moreover, we also point out The study also points out some current model limitations, here we particularly focus on implications with respect to the discrete representation of topography on a particularly the implications of representing topography and complex heterogeneous facades on a discrete Cartesian grid, complex heterogeneous facades, as well as and glass facades that are not well fully represented in terms of radiative processes. With these findings presented, we aim-
- 30 <u>Our findings are able</u> to validate the representation of physical processes in PALMas well as to point, while pointing out specific shortcomings. This will help to build a baseline for future developments of the model and for-improvements of simulations of physical processes in an urban environment.

1 Introduction

A large percentage majority of the world's population live in large cities (55 % as of 2018) and the this percentage is expected

- to be growing (UN, 2019). In addition to thatgrow (UN, 2019). At the same time, global climate change, especially global temperature increase, will influence most natural ecosystems increases, will be influencing nearly every natural ecosystem and human society, with potentially severe impacts worldwide. The high level of attention currently being paid to the impact of climate change on urban areas is eertainly legitimate and therefore amply justified, and is supported by many important studies and globally adopted reports (IPCC, 2014a, b). Moreover, intensified urbanization raises reports of global standing
- 40 (IPCC, 2014a, b). This intensifying urbanization heightens the awareness that control of the microclimate in the urban environmentis crucial for well-being of city inhabitants, as it, which can reduce heat stress and contribute to improvements of the living environment in cities. among other general environmental improvements, is crucial for the well-being of city inhabitants (Mutani and Fiermonte, 2017). The problem of increased heat stress in urban areas as a consequence of what has become known as the urban heat island (UHI) is therefore of direct concern to the municipal authorities municipal authorities.
- 45 who are well aware that the physical well-being of their inhabitants is vital , in many ways, to the well-being of the whole city. Moreover, the UHI effect is often followed by secondary processes, e.g. such as air quality issues. Researchers have responded to, or anticipated, such concern about and the requirement for modelling of urban climate processes and various, and several small-grid scale models and frameworks for (numerical) numerical climate modelling have recently been developed (Geletič et al., 2018).
- 50 The health and well-being of the urban population is influenced by the conditions of the urban environment. The local microclimate, exposure to pollutants, as well as human comfort of the urban population and general human comfort depends strongly on the local conditions determined driven by the urban environment. Thereby, the The turbulent flow, exchange of latent and sensible heat, as well as the and radiative transfer processes play an important role in the urban microclimate and need to be considered in modelling approaches. Implementation The implementation of important microclimate processes
- 55 (e.g. turbulence, heat fluxes or and radiation) in street-level scale models is typically partially or fully parameterized. The most exhaustive approach consists of a group of computational fluid dynamics (CFD) models. The explicit simulation of turbulent

flow is computationally demanding; thus, various techniques have to be adapted to make calculations feasible, usually based on limiting the range of the length scales and time scales of the turbulent flow to be resolved.

This study is using uses the PALM model system 6.0 (Maronga et al., 2020), which is an atmospheric modelling system.

- 60 The core of the system contains model dynamics based on the LES (Large Eddy Simulation) and RANS (Reynolds-Averaged Navier-Stokes) techniques with additional modules for modelling of various atmospheric processes, e.g. interaction of atmosphere with earth surface or cloud microphysics. This system core is complemented with a rich set of *PALM-4U* (PALM for urban applications) modules related to modelling of physical phenomena relevant for urban climate, such as the interaction of solar radiation with urban surfaces and with urban vegetation, sensible and latent heat fluxes from the surfaces, storage of
- 65 heat inside buildings and in pavements, or dispersion and chemical reaction of air pollutants (see Maronga et al., 2020). The first version of the PALM urban components represented the urban surface model (PALM-USM) which had has been validated using data from a short experimental campaign in the centre of Prague (Resler et al., 2017). The new set of modules in PALM is more general and is divided according to the physical processes they cover. The most relevant for urban climate are the land surface model (LSM), the building surface model (BSM), the radiative transfer model (RTM), and the plant-canopy
- 70 model (PCM), and the chemistry transport model (CHEM). The human biometeorology module (BIO) then allows to evaluate evaluation of the impact of simulated climate conditions on the human population.

Validation of the urban model requires a dataset of measurements of the urban meteorological and air quality conditions, the properties of the urban canopy elements and of and the energy exchange among parts of the urban canopy. Several campaigns of comprehensive observations and measurements of the urban atmospheric boundary layer, covering more than one

- 75 seasonwere performed, have been done in the past. The Basel Urban Boundary Layer Experiment (BUBBLE) dataset containing observations from Basel is specifically targeted for validation of urban radiation models, urban energy balance models, and urban canopy parameterizations (Rotach et al., 2005). The MUSE experiment (Montreal Urban Snow Experiment) aimed at the thermoradiative exchanges and the effect of snow cover in the urban atmospheric boundary layer (Lemonsu et al., 2008). The CAPITOUL (Canopy and Aerosol Particles Interaction in TOulouse Urban Layer) project (Masson et al., 2008) concentrated
- 80 on the role of aerosol particles in the urban layer.

Results of urban measurement campaigns have already been used for validation of several micrometeorological models, models of radiative transfer, and microscale chemical transport models. Micro-scale model validation brings difficulties due to high heterogeneity of the urban environment and the studied quantities, due to modelled variables, uncertainty in the knowledge of the details of detailed knowledge of urban canopy properties, as well as due to local irregularities caused by domain dis-

- 85 cretization. Important examples of such validation studies were published by Qu et al. (2013), Maggiotto et al. (2014) or Toparlar et al. (2015). Most often they and Toparlar et al. (2015). These validation studies most frequently analyze micrometeorological models of the RANS type. Early examples of LES validation studies that include thermal conditions within cities were presented by Nozu et al. (2008) and Liu et al. (2012). Due to our previous experience with a limited validation of surface temperatures simulated by the PALM model (Resler et al., 2017), the aim of this study was to design a comprehensive
- 90 experiment for complex model validation, including air velocity, air pollutionor, and surface temperature analysis.

The main goal of this study was to compare detailed temporary. The focus on collection of detailed temporally and spatially localized observations with the results of the micro-scale PALM simulations in various urban canopy and meteorological conditions was dictated by the intention to use these observations to assess the performance of the newly developed or updated PALM modules RTM, BSM, LSM, and PCMinside the complete modelling system. Additional purpose of this study was to

- 95 assess the PALM model performance for its utilization in urbanistic studies. The results of this study serve for planning future improvements PCM, and CHEM. This focus of the study also complied with its additional purpose, which was assessment of the utility of the PALM model and its PALM-4U components as well as for improvements of the model inputs. These results also provide information which can improve the design of future validation campaignsperformance for detailed urban studies (Geletič et al., 2021).
- 100 These considerations also influenced the selection of the studied area. study area. The Prague-Dejvice quarter provides a typical urbanized area inside is an urbanized area typical of others in Prague and similar Central European cities with various types of urban environment. Further, the realization of the street level observation campaign was technically and organizationally easier in this area than e.g. in areas such as the historical centre of Prague. Moreover, this area represents one of the pilot areas for urban adaptations studies carried out in cooperation with Prague Municipality and their organizations
- 105 (e.g. Prague Institute of Urban Planning and Development). The selection of this area was thus also influenced by their Their interest in the results of this study and their plan of plans for subsequent modelling studies of urban heat island and air quality adaptation and mitigation strategies for this quarter also influenced our selection of this area.

Section 2 gives a detailed overview of the observation campaign, followed by a description and an evaluation of the numerical setup in Sect. 3 and 3.5. In Sect. 4 results from the numerical experiment and the observation campaign are presented and 10 compared. Finally, Sect. 5 closes with a summary, outlines the current limitations of the model, and gives ideas for future improvements.

2 Observation campaign

The observation campaign was designed with two main aims: 1) to evaluate PALM's capability, with its newly developed or improved thermal modules - capability from the radiative transfer model (RTM), land and building surface modules (LSM,
BSM), and plant canopy model (PCM)through its capability, to reproduce surface temperatures; 2) to evaluate its capability to reproduce pollutant concentrations and meteorology values meteorological quantities in different types of street canyons, with special focus on the impact of trees in the located in streets on both types of quantities. The campaign was carried out in a warm part of the year (10–23 July 2018, further referred to as summer campaign) and a cold part of the year (23 November–10 December 2018, further referred to as winter campaign). Measurement locations are shown in Fig. 1 and measurements itself

120 the measurements themselves are described in Sect. 2.3.1–2.3.5. More details on the campaign are available in ČHMÚ (2020).

2.1 Study area

Study The study area is located in the north-west centre of Prague, capital city of the Czech Republic. The localization and position and a map of this area is are presented in Fig. S1 in the supplements. This figure also marks the extent of the PALM modelling domains; for more information about model domain setup see Sect. 3.1. The studied study area includes complex

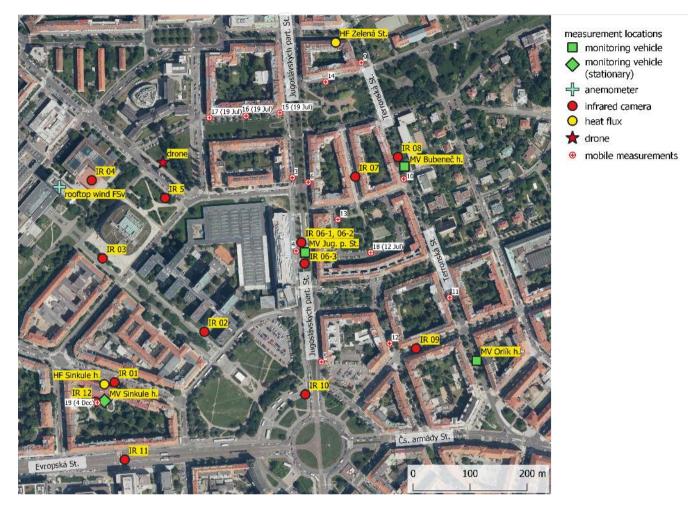
- 125 terrain mainly in the northern-western part of the outer domain , the altitudes range from up to (further referred as the parent domain), with altitude ranging from 175 to 346 meters above sea levelin the outer domain, . The altitude variability of the inner domain (further referred as the child domain) is up to 30 m (see Fig. S2). The observations were located inside the inner child domain (blue square in Fig. S2). This area is characterized as is a densely built-up area with specific conditions created by the roundabout (Vítězné náměstí) in combination with south-east-west-east (Evropská / Čs, Armády) and south-north-armády)
- 130 and north-south (Jugoslávských partyzánů / Svatovítská) oriented boulevards. East and partially south part of this domain represents. The eastern and southern parts of the child domain represent a typical historical residential area in Prague-Dejvice, with a combination of old and new buildings and a variety of other urban components, such as gardens, parks or parking places). North-west quarter is built up by and parking places. The north-west quarter has the larger buildings of the Czech Technical University campus. South-west and north-east The south-western and north-eastern parts of the domain are more
- 135 sparsely built-up by family houses. Local specific specific features include green intra-blocks with gardens and trees, usually with pervious ground surfaces; Prague historic centre usually has impervious intra-blocks. The building heights alongside the streets range approximately from 20 to 30 m, with the highest building in the domain is being 60 mhigh. Both boulevards are approximately 40 m wide and do not contain much contain little green vegetation, except for Jugoslávských partyzánů Street where high broadleaf trees (±which has some broadleaf trees about 20 m) are located. Majority high. The majority of the
- 140 trees are located in the intra-blocks and parks. Landcover The landcover map of the study area, based on Urban Atlas 2012 geodatabase, is shown in Fig. S3.

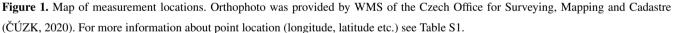
2.2 Validation episodes and synoptic situation

2.2.1 Summer campaign

The summer observation campaign ran for two weeks from 10 July 2018 to 23 July 2018 (see Table S2 in supplements), out of which two shorter episodes were selected for model simulations: 14–16 July (e1) and 19–23 July (e2). Synoptically, for most of the summer campaign the weather was influenced by a high pressure ridge over Central Europe between an Icelandic low and an Eastern Europe European low-pressure systemfor most of the summer campaign. Daily maximum temperature as measured at the Praha-Karlov (WMO ID 11519) station was below 30°C for the entire period, with the exception of 21 July

when the maximum temperature reached 31.2°C. The beginning of the period was partially cloudy, mostly with altostratus
clouds forming which formed in the morning and early afternoon on 19 July. The period between afternoon on the afternoon of 19 July and late afternoon on 21 July was mostly clear with cirrus clouds. End of the 21 July The end of the campaign was cloudy, mostly with low-level cumulus. Important solar parameters mid-episode (19 July 2018) were: time of sunrise at 03:13 UTC, time of sunset at 19:02 UTC and solar noon at 11:08 UTC.





2.2.2 Winter campaign

- 155 The winter part of the observation campaign lasted from 24 November 2018 to 10 December 2018 (see Table S3 in supplements) and for the purposes of model validation, three episodes were selected: 24–26 November (e1), 27–29 November (e2), and 4–6 December (e3). Weather was influenced by a typical late autumn synoptical situation with westerly flow and low-pressure systems and a series of fronts separated by two anticyclonic situations (27–29 November and 5 December). During the campaign several occluded frontal passages were recorded in Prague: 24, and 30 November and 2, 3, 4 and 6 December, with rainfall on 30 November (4.3 mm at Praha-Ruzyně station; WMO ID 11518) and 2 and 3 December (9.8 mm and 3.6 mm at Praha-Ryzuně
- Praha-Ryzyně station). Average daily temperatures ranged from -4 °C on 29 November to 9 °C on 3 December $\frac{2018}{2018}$. Average daily wind speed was around $3 \text{ m.s}^{-1} \text{ m s}^{-1}$, except for 26 November when it reached 4.4 m.s^{-1} and then m s⁻¹ and 4–6 Decem-

ber with daily values of 4.8, 6.0 and 5.7 m.s⁻¹. Important parameters of the solar radiation daily cycle m s⁻¹. Important diurnal solar radiation parameters in Prague were (as of 1 December 2018): sunrise at 6:39 UTC, sunset at 15:02 UTC, solar noon at 10:51 UTC.

165 10:51 UTC.

2.3 Observed quantities and equipment used

2.3.1 Infra-red Infrared camera measurements

Surface temperature measurements by the an infrared (IR) camera were carried out for during two days (45 hours) total) of the summer and 3 days (50 hours) during the summer and winter campaign respectively total) of the winter campaigns (see Ta-

- 170 ble S2 and Table S3). Measurements were performed taken at twelve locations shown in Fig. 1 approximately every 60–80 minutes. At each location, several directions were chosen and usually two snapshots capturing horizontal (groundsground) and vertical (wallswall) surfaces were taken in each direction. We use the following nomenclature further in the text: <location_number>-<direction_number>_H/V. For example 02-1_H means image of the ground taken from the second location in the first direction. In every image, a few evaluation points labeled labelled by numbers were chosen and temperature time series extracted. The
- 175 particular point at which modelled and observed values are compared is then referred to e.g. as 02-1_H3. The observation campaign in total gathered time series of surface temperature for 66 ground and 73 wall evaluation points, representing various surface types, in order to evaluate model performance under different surface parameter settings (e.g. such as different surface materials and conditions).

Temperature was measured by the FLIR SC660 (FLIR, 2008) - the same camera that was used in Resler et al. (2017). As

180 already described in the cited in this article, the camera's thermal sensor has a field of view of is 24 by 18° and a spatial resolution (given as an instantaneous field of view) of is 0.65 mrad. The spectral range of the camera is 7.5 to 13.0 μm, and the declared thermal sensitivity at 30°C is 45 mK. The measurement accuracy for an object with a temperature between 5 and 120°C, and given an ambient air temperature between 9 and 35°C, is ±1°C, or ±1% of the reading. The camera offers a built-in emissivity-correction option, which was not used for this study. Apart from the infrared pictures, the camera allowed us to simultaneously take pictures in the visible spectrum simultaneously.

Where possible, pictures were processed semi-automatically as described in Resler et al. (2017). This processing required requires the presence of 4 well-defined points occurring at in each picture, which were are used to correct for changes in camera positioning between the measurements, measurements as the camera was carried from one location to another. Pictures, rotated around locations. Pictures which did not allow for semi-automatic processing (mostly ground images) were handled manually and temperatures were extracted by the FLIR Tools v5.13.18031.2002 software (https://www.flir.com/products/flir-tools/). Ex-

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amples of semi-automatic and manually processed images are shown in Fig. S4.

Surface temperature measured by the FLIR SC660 was compared with the data from heat flux measurements at Sinkule house described in captured by the heat flux measuring system TRSYS01 (see Sect. 2.3.2). The results are shown in Fig. S5. The IR camera generally gives higher values than the TRSYS01 system (instantaneous measurements are compared with 10-min averages): in summer ground floor temperatures are on average 1°C higher (range of differences was between difference range 0.0and 2.8 - 2.8°C) and 1st floor on average 0.1°C higher (range of differences between -2.0 and +1.3°C). In winter the ground floor temperatures are on average 2.1°C higher (range of differences between difference range 0.5and 3.5 - 3.5°C) and 1st floor on average 1°C higher (range of differences between -0.6 and +2.0°C).

2.3.2 Wall heat fluxes measurement

- Heat fluxes through the building facade and window windows were measured by the high-accuracy building thermal resistance measuring system TRSYS01 equipped with two HFP01 heat flux plates and two pairs of thermocouples (TC). Operating-The operating temperature range of HFP01 and TC is -30 to +70°C. Declared The declared sensitivity of temperature difference measurements between inner and outer side sides of the wall is 0.02°C and heat flux measurement resolution 0.02W.m⁻². Calibration W m⁻². The calibration uncertainty of HFP01 is ±3 % (Hukseflux, 2020). Heat fluxes were measured through the north-east-facing wall of the Sinkule house and through the north-facing wall and window of the building in Zelená Street (Fig. 2). Position The position of the sensors on both buildings is shown in Fig. S6. Silicone glue was used to attach the sensors
 - to the outside wall in on the 1st floor of Sinkule house during the winter campaign. Otherwise sensors were mounted by a two-sided carpet tape.



Figure 2. Detail of heat flux sensors sensor and thermocouples thermocouple mounting. Left - Sinkule house 1st floor; center-left - Sinkule house ground floor; center-right - Sinkule ground floor - inner temperature sensor; right - Zelená Street. For Sinkule house and Zelená Street location locations see Fig. 1.

The Sinkule house was built before World War II with walls made of construction blocks. The ground floor wall is 34 cm
 thick , without insulation without insulation, and the facade is made of ceramic tiles. The wall in of the 1st floor is 41 cm thick including 6 cm thick polystyrene insulation on the outer side. Facade The facade surface is scratched plaster with scratches of 1–2 mm depth (see Fig. 2).

The house in Zelená Street is a typical representative of buildings in the area, with walls made of bricks construction blocks.

Wall thickness at the place of measurement was approx. 30 cm with 2.5-cm lime-cement plaster on the inner and outer side
 sides of the wall. Heat flux measurement through the window was not used in PALM validation and therefore is not described here.

A quality check measurement was done at the beginning of the summer campaign – sensors were placed side-by-side in the 1st floor of Sinkule house between 19 July 17:40 CEST and 20 July 12:00 CEST. Absolute The absolute difference of the facade surface temperature was 0.0–1.5°C with a median value of 0.1°C. Absolute The absolute difference of measured heat fluxes was 0.0–2.1 W.m⁻² W m⁻² with a median value of 0.6 W.m⁻² W m⁻².

2.3.3 Vehicle observations

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Air quality and meteorological measurements in the street canyons were obtained by two monitoring vehicles, which were shuttled periodically among the three locations marked as green squares in Fig. 1. One location was in Jugoslávských partyzánů Street (Jug. p. St.), an approx. 42-m wide boulevard with sparse trees. The two remaining locations were in the 25-m wide Terronská Streetnext to the, one next to Bubeneč house and the other next to Orlík house. Next to the Near to Bubeneč house , there are full-grown broadleaf trees with crowns covering the whole street. Broadleaf trees next to the near to Orlík house are smaller and their crowns are covering $\frac{2}{3}$ cover a maximum of two thirds of the street canyonin maximum. Buildings in all locations are approx. 25 m high. Pictures of the measurement locations are shown in Fig. S7. The observations were organised in a way so as to provide information about air quality and meteorological conditions in the three locations but and also to compare the east and west side sides of the street canyoncanyons. Each monitoring vehicle remained at a particular location

- for at least two whole days (see Table S2 and Table S3). Based on our own traffic census from 4–6 December 2018, the total workday load in Terronská St. (past Bubeneč house) is 7,700 vehicles, which is approximately 44 % of the traffic intensities in Jug. p. St. The number of small trucks (60) in Terronská St. is only 20 % of their count that in Jug. p. St. and the number of busses buses (20) is only 2 % of their count the number in Jug. p. St. There was only one large truck per day registered noted
- 235 in Terronská St., compared to approx. 80 in Jug. p. St. Apart from the street canyon measurements, one stationary monitoring vehicle was located in the yard of the Sinkule house during courtyard of Sinkule house throughout the whole campaign to provide the urban background meteorological and air quality values.

The vehicles in the street canyons were equipped with analyzers of NO_X, NO₂, NO, O₃, SO₂, CO, PM₁₀, PM_{2.5}, and PM₁ measured at the top of the vehicle roof (approx. 4.6 m). Calibrations of all air quality analyzers were performed during the

- 240 transfer between the transfer between locations to eliminate loss of data during parallel measurements. Meteorological variables measured included wind speed and direction and turbulent flow characteristics measured by the METEK 3D ultrasonic anemometer on a meteorological mast at a height of about 6.8 m above the ground (to reach above tree crowns fit under the tree crowns in Terronská St. next to Bubeneč house), air temperature, relative humidity, global radiation, and atmospheric pressure. One-minute averages of all entities were available from the instruments Wind and turbulent flow characteristics measured by
- 245 the METEK anemometer had a 10 minute resolution, with the remaining variables recorded by the instruments at 1-minute resolution. For further analysis and PALM evaluation, 10-minute averages of measured variables were used. Both vehicles

also had a video camera placed at the front windscreen. The recording was These recordings were then used for detailed time disaggregation of traffic emissions in the measured at the measurement location and for calibration of an automatic counting system (see Sect. 3.4).

250 The vehicle in the yard of the Sinkule house was measuring Sinkule house courtyard measured the same variables with the same time resolution except for the following differences: <u>PM1 was PM1, PM25</u> and turbulence characteristics were not measured; wind speed and direction were measured by the GILL 2D WindSonic anemometer at the standard height of 10 m.

2.3.4 Mobile measurements

On selected days of the measurement campaigns, mobile measurements using a dedicated monitoring vehicle were performed

- 255 to get more detailed information on air quality in the inner domainchild domain, mobile measurements using a dedicated monitoring vehicle were made (12, 18, 19 July, 26 November, and 4 December). The vehicle was moving. This vehicle travelled between the locations shown on Fig. 1, stopping and measuring in at each of them for five minutes. Two loops were made on every measurement day. On 19 July only one loop among locations 3, 6, and 15–17 was made, but the measurements took-with measurements taken over 15–20 min. The vehicle was equipped with NO_X, NO₂, NO, O₃, SO₂, CO, PM₁₀, PM_{2.5}, and PM₁
- 260 analyzers. A-Starting from the second measurement on 17 July, a Garni 835 weather station was used for an indicative measurement of temperature, wind and relative humiditystarting from the second measurement on 17 July 2018. Some measurements were not available on particular days details are given in Table S2 and Table S3.

2.3.5 Higher level observations

To get information about higher levels, two additional observations were proposed in the scope of the observation campaign used two other measurement platforms. The first one was a stationary measurement of wind flow above rooftops in the area of interest throughout the campaign duration on the top of the highest building in the child domain (approx. 60 m high). A 2D anemometer was installed on the flat roof of the Faculty of Civil Engineering of the Czech Technical University – the highest building in the inner domain (approx. 60 m high) (FSv; see Fig. 1). The anemometer was positioned approximately in the middle of the highest roof section, 2 m above the flat roof top. The location was the same in summer and winter campaigns. Measurement frequency was 1 second . The and 10-minute averages were used for further evaluation. The second one-was a measurement of vertical profiles in the lowest part of the atmosphere by drone. Originally, two one-day drone observation campaigns were scheduled. Due to administrative restrictions, the summer drone observations were not realised

and the winter part ones had to be moved from the center of inner centre of the child domain to the location marked in the Fig. 1. Also Additionally, the maximum flight altitude had to be limited to 80 m above ground. The drone was equipped with

275 the GRIMM Portable Laser Aerosol spectrometer and Dust Monitor Model 1.108 and a HC2A-S probe from ROTRONIC for the measurement of temperature and relative humidity measurements (ROTRONIC, 2020). Unfortunately, the probe showed a longer than expected relaxation time and was which meant the observation instruments were not able to adapt stabilize quickly enough during the descent. Recalculation of particle counts to mass concentration was also burdened with large errors. The obtained results results obtained were not reliable enough to be included in used for PALM validation, but temperature and

280 relative humidity profiles are provided in supplements (Fig. S8 and Fig. S9).

2.3.6 Standard CHMI observations used for validation

The following <u>Relevant</u> standard meteorological and air quality measurements were used for the evaluation of WRF and CAMx simulations (which provided initial and boundary conditions for PALM ; see as described in Sect. 3.3). This evaluation is presented in Sect. 3.6.13.5. WRF vertical profiles were evaluated against the upper air soundings from Praha-Libuš (WMO

ID 11520) station located in the south a southern suburb of Prague, 11 km apart from from the center of PALM inner domain. Radiosonde the PALM child domain. A radiosonde is released every day at 0, 6, and 12 UTC. For the evaluation of global radiation, two meteorological stations were selected: the already mentioned Praha-Libušstation and , and the Praha-Karlov (WMO ID 11519) station situated in a densely built-up area in the wider nearer the center of Prague approximately 4 km from the PALM inner child domain. PM₁₀ and NO_X concentrations provided by from the CAMx model were compared with the measurement measurements from automated air quality monitoring stations. Only the 5 background stations closest to the PALM inner child domain were used. Station locations are shown in Fig. S10. More detailed information about the stations is

given in Table S4 and Table S5.

Observations from the Praha-Ruzyně station (WMO ID 11518) situated at Prague airport approximately 9 km west from the center of PALM domain were used to evaluate WRF wind speed and, in conjunction with the campaign wind measurements on the FSv building roof, the modification of wind speed by the orography and buildings and how PALM captures this effect.

3 Model simulation setup

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3.1 PALM model and domains configuration

The PALM model system version 6.0 revision 4508 (Maronga et al., 2015, 2020) was utilized for this validation study. It consists of the PALM model core and embedded modules and of PALM-4U components which have been specifically devel-300 oped for modelling urban environments. The PALM model core solves the incompressible, filtered, Boussinesq-approximated Navier-Stokes equations for wind (u, v, w) and scalar quantities (potential temperature, water vapor mixing ratio, passive scalar) on a staggered Cartesian grid. The sub-grid scale terms that arise from filtering are parametrized using a 1.5-closure by Deardorff (1980) with modifications after Moeng and Wyngaard (1988) and Saiki et al. (2000). Buildings and orography are mapped onto the Cartesian grid using the mask method (Briscolini and Santangelo, 1989), where a grid cell is either 100% fluid

305 or 100% obstacle. The advection terms are discretized by a 5th according to fifth-order scheme after Wicker and Skamarock (2002). For temporal discretization, a 3rd-order third-order low-storage Runge-Kutta scheme (Williamson, 1980) is applied. The Poisson equation is solved by using a multigrid scheme (Maronga et al., 2015).

The following are the urban canopy related PALM and PALM-4U modules were modules employed in this study: the . The land surface model (LSM, Gehrke et al., 2020; to be submitted to GMD) was utilized to solve the energy balance

- 310 over pavements, water- and other natural-like surfaces, the water, and low-vegetated surfaces. The building surface model (BSM, formerly USM, see Resler et al., 2017 in previous versions and in Resler et al., 2017 called USM) was used to solve the energy balance of building surfaces (walls and roofs). The BSM was configured to utilize an integrated support for modelling of fractional surfaces (Maronga et al., 2020). Dynamic and thermodynamic processes caused by resolved trees and shrubs were managed by the embedded plant-canopy model (PCM). Radiation interaction between resolved scale vegetation, land-
- 315 surface, and building surfaces was modelled via the radiative transfer model (RTM, Krč et al., 2020; to be submitted to GMD). Downwelling shortwave and longwave (SW) and longwave (LW) radiation from the upper parts of the atmosphere, which were used as boundary conditions for the RTM, were explicitly prescribed from the stand-alone Weather Research and Forecasting model (WRF; see Sect. 3.3 for details) simulation output for the respective daysrather than, rather than being modelled by e.g. the Rapid Radiation Transfer Model for Global Models (RRTMG). This way, effects of mid- and high-altitude clouds on
- 320 the radiation balance were considered in the simulations. It is needed important to note that without applying the by not using RRTMG some physical processes were missed, such as vertical divergence of radiation fluxes leading to heating / cooling of the air column itselfwere missed, which; these may become especially important at nighttime. However, sensitivity tests with RRTMG applied revealed that the effect on nighttime air temperature was negligible in our simulations. In addition to the meteorological componentquantities, the embedded online chemistry model (Khan et al., 2020; to be submitted to GMD) was
- 325 applied to model concentrations of NO_X, PM₁₀, and PM_{2.5}. Chemical reactions were omitted in this case to simulate purely passive transport of the pollutants. For a human thermal comfort estimatimation the PALM biometeorological module (?) was used. However, campaigns were not designed for their evaluation and these results were not validated.

Additionally both self-Both self and online nesting features of PALM-4U PALM were utilised. Self-nesting means that a domain with a finer resolution can be defined inside a larger domain and this subdomain (child domain) receives its boundary
conditions from the coarse-resolution parent domain at every model timestep. Here, a one-way nesting without any feedback of the child simulation on the parent simulation (Hellsten et al., 2020; to be submitted to GMD) was applied. The coarse-resolution parent simulation itself received its initial as well as lateral and top boundary conditions from the simulations of the mesoscale model WRF transformed to a PALM dynamic driver (see Sect. 3.3). The boundary values were updated at every model time step, and this This process is hereafter referred to as offline mesoscale nesting (Kadasch et al., 2020; to be submitted to GMD). As the offline nesting was used for coupling to). The values of the velocity components, potential temperature, and values for the mixing ratio at the lateral and top boundary were updated at every model time step, while

- linear interpolation in time was used to interpolate between two WRF timesteps. The WRF solution was mapped fully onto the boundaries starting at the first grid point above the surface; boundary grid points that lie below the surface were masked and were not considered further. As the mesoscale model that does not resolve turbulence, the turbulence was triggered at the
- 340 model boundaries using an embedded synthetic turbulence generator (STG) according to Xie and Castro (2008), which imposed spatially and temporally correlated perturbations every time-step onto the velocity components at the lateral boundaries. For additional details on PALM's mesoscale nesting approach we refer to Kadasch et al., 2020.

The initial and boundary concentrations of modelled pollutants of the parent domain were taken from simulations of the CAMx model (Comprehensive Air-quality Model with Extensions; see Sect. 3.3). For more detailed information about the

345 PALM model, embedded modules, and the PALM-4U components see Maronga et al. (2020) and the companion papers in this special issue.

The location of the parent and child modelling domains is shown by in Fig. S1. The parent domain extends horizontally by $4 \times 4 \text{ km}^2$ in the x- and y-direction km² in the x- and y-direction, respectively, with an isotropic grid spacing of 10 m. The vertical z-direction z-direction is covered by 162 layers for summer and 82 layers for winter simulations, respectively.

- The vertical grid spacing is 10 m for the lower 250 m of the domain, <u>Above 250 m</u>, when the height was well above the building-affected layer, the vertical grid was successively stretched up to a maximum vertical grid spacing of 20 m in order to save computational resources. The domain top is at 2,930 m for summer and 1,330 m for winter simulations, respectively. This extent safely covers the convective layer with a sufficient buffer. We note that the 10 m resolution of the parent domain is sufficient to explicitly resolve the majority of the buildings and trees (see Fig. S11 and Fig. S12 in supplements) which means
- that no additional parameterization of the urban canopy is needed. The child domain extends by extent is $1,440 \times 1,440 \times 242 \text{ m}^3$ m³ in the x-, y-, and z-direction, z-directions respectively, with an isotropic grid spacing of 2 m.

Parent and child domains were initialized by vertical profiles of u, v, w, potential temperature and mixing ratio, and soil moisture and soil temperature, transformed from WRF simulations (see Sect. 3.3). Since the initial soil and wall temperatures from a mesoscale model are only a rough estimate due to its aggregated nature, the PALM spin-up mechanism was applied (Managa et al. 2020). During a 2 day arigment the struggebaria and are my table of and endy the LSM and DSM teacther

360 (Maronga et al., 2020). During a 2-day spinup, the atmospheric code was switched-off and only the LSM and BSM together with the radiation and RTM model were executed. By this <u>method</u>, the material temperatures were already <u>eloser_close</u> to their equilibrium value and significant changes in material temperatures at the beginning of the simulation were avoided.

3.2 Urban canopy properties

Data availability, their harmonizationand costs, and cost/efficiency trade-offs often needs need to be considered (Masson et al., 2020). For solving the energy balance equations as well as for radiation interactions, BSM, LSM, and RTM require using the use of detailed and precise input parameters describing the surface materials (e.g., such as albedo, emissivity, roughness length, thermal conductivity, thermal capacity, and capacity and thermal conductivity of the skin layer). Also the plant canopy (trees and shrubs) is important as it affects the flow dynamics, heating, and evapotranspiration as well as the radiative transfer within the urban environment. Urban and land surfaces and sub-surface materials become very heterogeneous in a real urban environment when going to very fine spatial resolution. Any bulk parameterization for the whole domain setting would therefore be inadequate. Instead, a detailed setting of these parameters was supplied everywhere wherever possible. To obtain the needed detailed data, a supplemental on-site data collection campaign was carried out and a detailed database of geospatial data was created. Land-cover data are based on a combination of national (ZABAGED) and city of Prague (Prague OpenData) databases. ZABAGED geodatabase (ČÚZK, 2020) distinguishes 128 categories of well-targeted geographical ob-

375 jects and fields(e.g., for example built-up areas, communications, hydrology, vegetation, and surface). The Prague OpenData geodatabase (Prague Geoportal, 2020) distinguishes many local, user-specified –GIS layers, e.g. plans with showing actual and future development, land-cover for architects, a photogrammetry-based digital elevation model (DEM) etc. Building heights were available from the Prague 3D model, maintained by the Prague Institute of Planning and Development. For the first tree

canopy data mapping, LiDAR scanning was used in combination with a photogrammetric-based DEM. Derived heights were

- 380 manually calibrated using <u>data from the</u> terrain mapping campaign and extended with additional parameters (e.g., like crown height, width and shape, <u>and</u> trunk height and width). All descriptions of surfaces and materials and their properties were collected in GIS formats and then preprocessed into <u>the a PALM NetCDF</u> input file corresponding to the PALM Input Data Standard (PIDS; Heldens et al., 2020). This file includes information on wall, ground, and roof materials and properties similar to <u>that those</u> used to estimate surface and material properties in Resler et al. (2017) and Belda et al. (2020).
- Each surface is described by material category, albedo, and emissivity, BSM surfaces additionally carry thickness, and window fraction. Parameters, such as thermal conductivity and capacity, are assigned to categories and estimated based on surface and storage material composition. In the case of walls and roofs, which are limited to four layers in the current version of BSM, this means the parameters of the two outer layers were assigned according to the properties of the covering material (e.g. plaster or insulation), while remaining layers were initialized by the properties of the wall material (e.g. bricks,
- 390 construction blocks, concrete, insulation). Wall and roof properties are described in table located in Table S6. In case of For pavements and other LSM surfaces, all parameters except albedo and emissivity were assigned according to the PALM LSM categories.

Each tree in the child domain was described detailed by its position, diameter, trunk parameters, and vertically stratified base leaf area density. The actual distribution of the leaf area density (LAD) within the treetop was then calculated according

395 to the available light exposure of the particular gridbox inside the treetop according to the Beer-Lambert law, leading to lower LAD in the centres of large and/or dense treetops. Note, at At the moment PALM does not consider the effect of trunks on the dynamic flow field and the thermodynamics; only LAD is considered. However, for the winter case leafless deciduous trees were considered by to be 10 % of their summer LAD to account for the effect of trunks and branches on the flow field.

3.3 Initial and boundary conditions

- 400 Initial and boundary meteorological conditions for the parent domain of the PALM simulations were obtained from the WRF model (Skamarock et al., 2008), version 4.0.3. The WRF model was run on three nested domains, with horizontal resolutions of 9 km, 3 km and 1 km and 49 vertical levels. The inner child domain has 84 × 84 grid points in the horizontal. The configuration was standard but parameterizations have been chosen choice of configuration started from the most usual settings for the given resolution and required latitude. Then minor variations in parameterizations were tested so as to decrease possible discrepancies
- 405 which might arise from boundary conditions . provide the best possible boundary conditions to PALM for each simulation. Consequently the NOAH LSM (Chen and Dudhia, 2001) and RRTMG radiation (Iacono et al., 2008) have been used in all simulations. As for PBL parameterization, Yonsei University scheme (Hong et al., 2006) has been Urban vs. non-urban parametrizations for PBL were tested and, as a result, the Yonsei University PBL scheme (Hong et al., 2006) was chosen for the summer episodes while for the winter episodes the Boulac urban PBL (Bougeault and Lacarrère, 1989) scheme has been
- 410 used. Except for thatgave a better agreement with observations. With this exception, no other urban parameterization has parameterizations have been used in the WRF model. MODIS land use categories have not been altered. The WRF output data have been collected from overlapping runs of length 12 hours, WRF was initialized from the GFS operational analyses and

predictions forecasts and output data from overlapping WRF 12 hour runs was collected. The first six hours of each run served as a spin-up. The boundary conditions for the offline nesting have been mesoscale nesting were then generated from forecast

415 horizons 7–12.

Air quality simulations that served as chemical initial and boundary conditions were conducted made using the chemistry transport model (CTM) CAMx version 6.50 (ENVIRON, 2018). CAMx is an Eulerian photochemical CTM that contains multiple gas phase chemistry options (CB5, CB6, SAPRC07TC). Here, the CB5 scheme (Yarwood et al., 2005) was invoked. Particle matter was treated using a static two-mode approach. Dry deposition was calculated following Zhang et al. (2003) and for wet deposition, the Seinfeld and Pandis (1998) method was used. To calculate the composition and phase state of the ammonia-sulfate-nitrate-chloride-sodium-water inorganic aerosol system in equilibrium with gas phase precursors, the

ISORROPIA thermodynamic equilibrium model was used (Nenes et al., 1998). Finally, secondary organic aerosol (SOA)

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chemistry was solved using the semi-volatile equilibrium scheme SOAP (Strader et al., 1999).

CAMx was coupled offline to WRF, meaning that CAMx ran upon on WRF meteorological outputs. WRF outputs were
then translated to CAMx input fields using the WRFCAMx preprocessor provided along with the CAMx source code (see http://www.camx.com/download/support-software.aspx). For those CAMx input variables that were not available directly in WRF output, diagnostic methods were applied. One of the most important inputs for CAMxthat, which drives the vertical transport of pollutants, is the coefficient of vertical turbulent diffusion (Kv). Kv is a significant parameter that determines the city scale air pollution and it is substantially perturbed by the urban canopy effects (Huszar et al., 2018a, b, 2020a, b). Here, the "CMAQ" scheme (Byun, 1999) was applied for Kv calculations.

- WRF and CAMx outputs were then postprocessed to into the PALM dynamic and chemistry driver. The data were transformed between coordinate systems and a horizontal and vertical interpolation was applied. As the coarse-resolution model terrain would not match the PALM model terrain exactly, the vertical interpolation method included terrain matching and the atmospheric column above the terrain was gradually stretched following the WRF hybrid vertical levels as they were converted
- 435 to the fixed vertical coordinates of the PALM model. The interpolated airflow was adjusted to enforce the mass conservation. Detailed technical description of the 3D data conversion procedure is beyond the scope of this studygiven in the supplements in Sect. S6. The Python code used for processing the WRF and CAMx data into the PALM dynamic driver file has been included in the official PALM distribution and published in the PALM SVN repository since revision 4766 in the directory trunk/UTIL/WRF_interface.
- Emission data for Prague used in the CAMx model were the same as described in the next following chapter. Other emission inputs are described in detail in Ďoubalová et al. (2020).

3.4 Emission data

Air pollution sources for our particular case are dominated by the local road traffic. Annual emissions totals were based on the traffic census 2016 conducted by the Technical Administration of Roads of the City of Prague – Department of Transportation

445 Engineering (TSK-ÚDI). Emissions itself The emissions themselves were prepared by ATEM (Studio of ecological models; http://www.atem.cz) using the MEFA 13 model. Jugoslávských partyzánů and Terronská StreetStreets, where air quality was measured during the campaigns, were <u>both</u> covered by this census. Emissions from streets not included in the census were available <u>in on a</u> grid with 500m spatial resolution. These emissions were distributed between the streets not covered by the census according to their parameters. Particulate matter (PM) emissions included resuspension of dust from the road surface

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(Fig. 3). Time disaggregation was calculated using a Prague transportation yearbook (TSK-ÚDI, 2018), public bus timetables, and our own short-time census (19–21 July and 4–6 December; days in which traffic intensities were derived from camera records). This time disaggregation was the same for the primary emissions (exhaust, brake wear etc.) as well as for resuspended dust. Higher dust resuspension caused by sprinkle material during winter time was not considered.

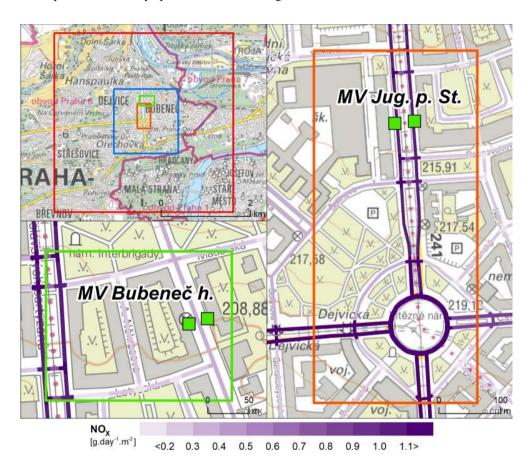


Figure 3. Nitrogen oxides (NO_x) emitted by the cars along their trajectories in selected locations in Prague-Dejvice. Concentrations Emissions were summarized in g.day⁻¹.m⁻² g day⁻¹ m⁻² and disaggregated to 1-hour time steps. Green The red and blue squares represent in the top left map indicate the extent of the parent and child PALM domain, respectively. The orange and green rectangles show the location of the expanded views at right and lower left. The expanded views show as green squares the air-quality measurement locations (MV) in Terronská St., Bubeneč house (lower leftbottom) and Jugoslávských partyzánů St. (right). The base map of the Czech Republic at 1:10,000 for the city of Prague was provided by Czech Office for Surveying, Mapping and Cadastre (ČÚZK, 2020).

Traffic data were supplemented by emissions from stationary sources from the Czech national inventory REZZO: point.

- 455 Point sources correspond to the year 2017(, the latest year available at the time of model input preparation). Residential heating was based on 2017 inventory and rescaled to 2018 by multiplying by the ratio of degree days DD(2018)/DD(2017); DD (r) is the sum of the differences between the reference indoor temperature and the average daily outdoor temperature on heating days. Residential heating emissions were available on elemental dwelling units urban areas with average area 0.5 km², km² and were spatially distributed to building addresses, where local heating source is registered, proportionally sources are registered.
- 460 in proportion to the number of flats. Time disaggregation of point source emissions was based on monthly, day-of-week, and hour-of-day factors (Builtjes et al., 2003; available also in Denier van der Gon et al., 2011). Residential heating emissions were distributed allocated to days according to the standardized load profile of natural gas supply for the households, which use it for heating only (Novák et al., 2019; OTE, 2020). Daily variation of residential heating emissions was taken from Builtjes et al. (2003).
- 465 All these input emission data were processed into the PALM input NetCDF files corresponding to the PALM Input Data Standard (PIDS).

3.5 Observation operator

To compare modelled and observed values, an observation operator which links model variables to observed quantities is needed. In case of For vehicle measurements, the situation was straightforward; horizontally, we used atmospheric quantities

- 470 and chemical compounds at the grid cell closest to the real placement of the sensors while vertically, we performed linear interpolation to the real height of the sensor. This approach was sufficient given the fine 2m resolution within the child domain. In case of For surface observations at grid-aligned surfaces (walls parts without wall sections without significant influence of step-like structures), the modelled values at the nearest grid face according to the real_actual placement of the sensor or evaluation point were also taken. However, at non grid-aligned walls, i.e. walls which are oriented in one of the south-west,
- 475 south-east, north-west, and north-east directions, walls are approximated by step-like structures and choosing the nearest grid face is not uniqueany more no longer unique, as illustrated in Fig. 4. In this ease these cases, the orientation of the real wall cannot be sufficiently represented by the one grid face but is approximated by grid faces with perpendicular orientation. HenceFor this reason, we virtually sampled surface quantities at both the two perpendicular surfaces and calculated the modelling counterpart of the observation as the average of these values. In the followinggraphs of the surface temperature, the sampled values
- 480 are plotted with suffixes "_l" and "_r" to distinguish between "left" and "right" adjacent faces and their average as by thin dashed lines in addition to their average representing the modelled value Implications which is shown by thick solid lines. Implications of this for the model evaluation as well as for the comparability of the model to the observations accompanied with the grid discretization are discussed in Sect. 4.1.7, along with the grid discretization.

Evaluation of the driving synoptic-scale model simulation setup]Results

485 Evaluation of the driving synoptic-scale model simulation setup To ensure the correct model couple setup and correspondence to general meteorological conditions, basic characteristics are evaluated in this section. This includes the evaluation of the

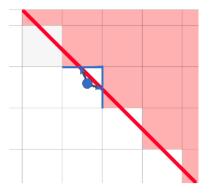


Figure 4. Illustration of Sketch to illustrate the mapping of a wall surface observation point on to a gridded step-wise approximation of the wallinto two. The red line represents the real wall surface, light grey lines delineate the grid cells, the light red area shows the footprint of the gridded building, blue circle the surface evaluation point, and the blue arrows represent the assignment of this point to the grid faces (blue lines) used for calculation of the corresponding modelled values.

driving synoptic-scale simulations of WRF and CAMx models, the vertical representation of the boundary layer in PALM, and the spatial development of the turbulent flow characteristics from the boundaries of the PALM parent and child domains. Special focus is put on the summer e2 and winter e3 episodes, in which IR camera observations took place. A description of the statistical methods used is given in the appendix A.

3.5.1 Meteorology

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3.6 Meteorology

3.6.1 Evaluation of the driving synoptic-scale simulation

Since the boundary conditions for the PALM simulations come from a model simulation as well, we need to check for potential misrepresentation of the real weather atmospheric conditions. First we assess the overall performance of the WRF model simulation on the synoptic scale by comparing the results with the known state of the atmosphere, represented here by the ERA-Interim reanalysis and atmospheric soundings obtained by the CHMI weather balloons radiosondes (downloaded from the University of Wyoming database; http://weather.uwyo.edu/upperair/sounding.html). Fig. ?? \$13 and Fig. \$14 show maps of geopotential height at 500 hPa and 850 hPa comparing the results of the WRF simulation (9km domain) with the ERA-Interim
reanalysis. The rest of the maps is in supplementary files; see Fig. \$11 and Fig. \$12.

Generally, the WRF simulations, being driven by the GFS, correspond very_Global Forecast System (GFS), correspond well to the ERA-Interim reanalysis in terms of the 500 hPa geopotential height field, with some shifts of the pressure field eastward on 19 July and northward on 21 July. Geopotential height at 850 hPa is also very well represented with some added detail, mainly during the day of the summer month in the summer due to a better resolved topography in the higher-resolution regional model simulation

505 model simulation.

Geopotential height at 500 hPa (left) and 850 hPa (right) for 20 July 2018 00:00 UTC. Blue line is ERA-Interim reanalysis, red line is WRF simulation of 9-km domain. Background layer uses the public domain shaded relief map from Shaded Relief (2020)

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Next, we evaluated Additionally, we compared the WRF results with atmospheric soundings for the station closest to our do-

- 510 main of interest, Praha-Libuš(note: the weather balloon, which is about 11 km south-southeast of the modelled area. Figures 5 and 6 show observed and modelled profiles of the potential temperature and wind speed at the sounding location for 20–21 July (episode summer e2) and 4–5 December (episode winter e3), respectively. Graphs for other episodes are provided in supplements (Fig. S15, S16, and S17). The radiosonde measurements are taken only-three times per day at 00, 06 and 12 UTC). Fig. ?? shows the vertical profile of potential temperature for 20–21 July and Fig. ?? for 4–5 December (vertical. The
- 515 modelled values are inferred from the 1 km resolution WRF model. In order to estimate spatial variability and consequently the utility of the sounding for validation of the WRF profiles within the PALM domain, WRF profiles for the rest of the episodes are included in the supplement; see Fig. S13, Fig. S14 and Fig. S15). Modelled profiles centre of the PALM domain are also shown. Modelled profiles from the PALM parent domain simulation are also included in these graphs; these are discussed in Sect. 3.7 below.
- 520 WRF profiles of potential temperature generally correspond well with the measurements observations with some notable exceptions in the surface layer (which is of near the surface, where WRF tends to underestimate nighttime stability and shows less marked near-surface instability during daytime in the summer case. However, here we emphasize that the near-surface profiles might also be affected by the fact that the relevant WRF model surface is not necessarily representative of local detail. The WRF wind-speed profiles also mainly reflect the conditions as observed, with a well modelled nighttime low-level jet (e.g.
- 525 21 July at 00:00 UTC, 5 December at 06:00 UTC). However, compared to potential temperature, modelled wind speed exhibits larger discrepancies to observations at various times, for example 20 July at 00:00 and 21 July at 12:00, and also tends to be higher, especially near the surface in the winter scenario. As discussed in the preceding paragraph, the radiosonde location is not within the PALM model domain, However, WRF profiles at the radiosonde location and the highest importance for this study, as the boundary conditions for the PALM simulations are taken from the lower levels) where the model tends to show a
- 530 lower diurnal range underestimating stability in the night time and instability during the day. Vertical profile of potential temperature from the soundings balloon observations at the Praha-Libuš station (magenta) and the nearest grid box of the WRF simulation 1km domain (evan) for 20–21 July.

Vertical profile of potential temperature from the soundings balloon observations at the Praha-Libuš station (magenta) and the nearest grid box of the WRF simulation 1km domain (cvan) for 4–5 December, PALM domain center show only

535 marginal differences. Hence, we are confident that the modelled boundary layer profiles from WRF, which are used as boundary conditions for PALM, are a sufficiently good representation of reality for this study.

Here we note that in case Another factor needing consideration is that in the summer cases the boundary-layer depth during the daytime is within the range of the 1 km horizontal grid resolution , which is especially the case for the summer simulations, Ching et al. (2014) and in the WRF simulations. Ching et al. (2014); Zhou et al. (2014) showed that in such situations resolved

540 scale convection can developwhich depends on the horizontal grid resolution as well as on the applied, also altering the

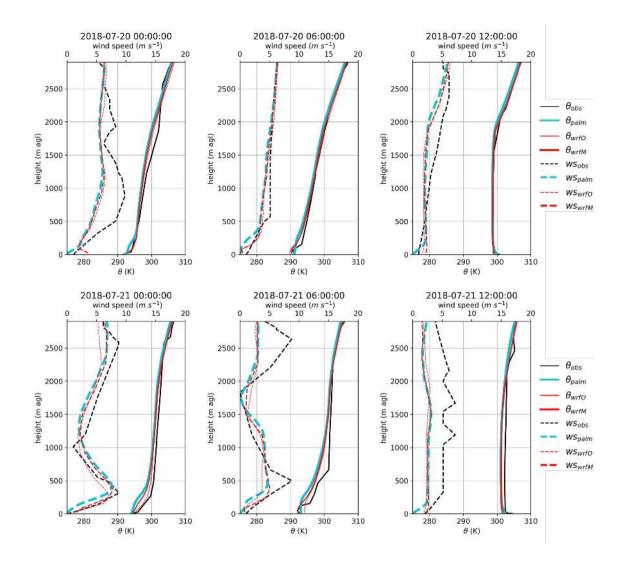


Figure 5. Vertical profiles of potential temperature and wind speed from the radiosonde observations at Praha-Libuš station for 20–21 July, with corresponding WRF (1 km horizontal resolution) and PALM (average from parent 10 m resolution domain) profiles. The potential temperature is represented by the solid lines, the wind speed is denoted by the dashed lines. The black line is the sounding observation, the cyan line the PALM model, the red line the WRF model. The thin red line is the WRF model at the sounding location while the thick red line is the WRF model in the centre of the PALM domain.

boundary-layer scheme. For a nested LES representation and leading to too large a vertical energy transport. For an LES nested into a mesoscale WRF simulation, Mazzaro et al. (2017) showed that such under-resolved convection may propagate into the LES domain, biasing the location of the up-updrafts and downdrafts. In order not to bias our simulation results by under-resolved convection in WRF propagating into the LES, we checked the WRF-simulation output for the occurrence of under-resolved convection but did not find any (not shown).

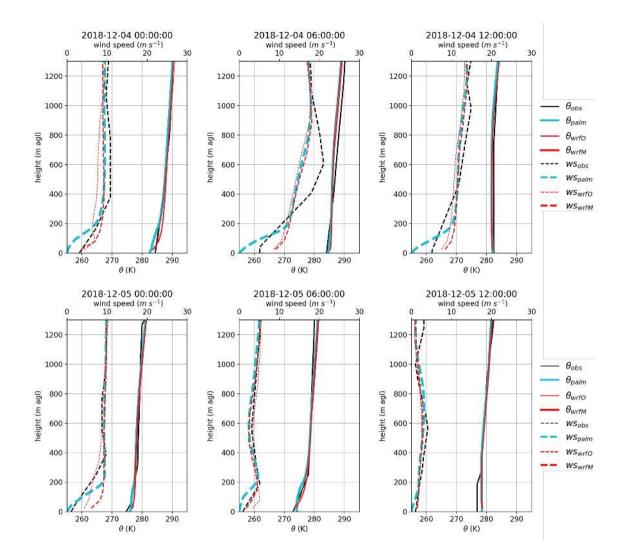


Figure 6. Vertical profile of potential temperature and wind speed from the radiosonde observations at Praha-Libuš station, with corresponding WRF (1km horizontal resolution) and PALM (average from parent 10 m resolution domain) profiles for 4–5 December. The potential temperature is represented by the solid lines, the wind speed is denoted by the dashed lines. The black line is the sounding observation, the cyan line the PALM model, the red line the WRF model. The thin red line is the WRF model at the sounding location while the thick red line is the WRF model in the centre of the PALM domain.

In our setup, PALM simulations need radiation values the PALM simulations we prescribed the incoming long and shortwave radiation obtained from the WRF simulationsas one of boundary conditions. To check for potential errors in these boundary conditions incoming radiation, we compare (see Fig. 7) global downwelling SW radiation as simulated by WRF (innermost lkm domain; black dots) in the grid box centered over the area of interest covering the center of the PALM child domain with observations at two CHMI stations in Prague with continuous global downward short wave radiation measurements: Praha-

Karlov (approx. 4 km southeast from the modelled area), and Praha-Libuš (approx. 11 km south-southeast from the modelled area(Fig. 7). WRF simulations show good agreement with observations in the summer campaign, with some overestimation of the global SW radiation on 14 and 23 July at noon, most likely due to which we attribute to the underestimation of cloud cover in the WRF simulation. During the winter campaign, the global downwelling SW radiation in WRF agrees with the observation

at nighttime as well as during the morning and afternoon hours. However, on 26, 28, and 29 November, and on 5 December, while WRF significantly overestimates the global radiation at noon SW radiation on other days due to underestimated cloud coverduring this period of time.

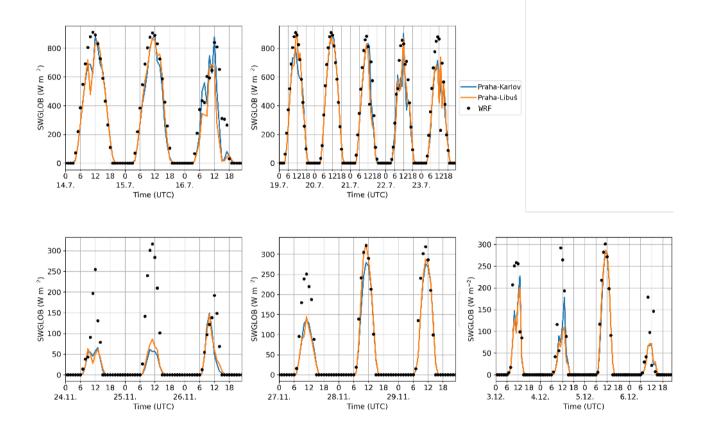


Figure 7. WRF modelled and observed <u>global downwelling SW</u> radiation for modelling episodes summer e1, e2 (top row) and winter e1, e2, e3 (bottom row)from : CHMI <u>station</u> Praha-Karlov (blue <u>line</u>), ; CHMI station Praha-Libuš (orange <u>line</u>); WRF simulation (black <u>dots</u>).

3.6.2 Air qualityBoundary-layer representation in PALM

For the CAMx model validation, urban background air quality monitoring stations closest to the PALM outer domain were 560 used (see Sect. 2.3.6). Validation was performed for hourly average concentrations of NO_X and PM₁₀. Only PALM episodes were included (i.e. 14–15 and 19–23 July for summer and 24–29 November and 4–6 December for winter). Metrics according to Britter and Schatzmann (2007) and Chang and Hanna (2004) for both campaigns are summarized in Table 1 (note that fractional bias values are positive, when model underpredicts observations). R statistical software (?) and openair package (Carslaw and Ropkins, 201 was used.

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For NO_X , the metrics show a significant underprediction of the measured concentrations (FB ca. 0.8) for both summer and winter episodes. Nevertheless, the daily variation is captured quite well (see Fig. S16) although in winter modelled peaks in the evening are larger then in the morning, while in the observed data it is just opposite. Analysis of PALM simulated concentrations in section 4.4 will show that this bias is not present there.

Summer PM₁₀ concentrations are less underestimated (FB ca. 0.5) and morning and evening peak is more sharp and appears

570 ca. 1 h earlier then in observations. Winter PM₁₀ are even a bit overestimated but CAMx model is not able to represent their real daily variation. Modelled daily variation is very similar to that for NO_X, which indicates that it is dominated by daily variation of traffic, while in reality different sources play more important role.

Evaluation of CAMx 1-h concentrations against urban background stations. NO_X NO_X PM₁₀ PM₁₀ Summer Winter Summer Winter N 684 816 907 1078mean obs μ g.m⁻³22.6 59.5 22.1 30.4mean mod μ g.m⁻³10.1 24.4 13.4 33.3FB 0.76 0.84 0.49 -0.09NMSE 1.51 2.15 0.65 0.53FAC2 0.38 0.31 0.50 0.69R 0.54 0.28 0.34 0.13

3.7 PALM simulation characteristics, vertical profiles

In order to check whether the observed boundary-layer structure is represented realistically by the LES simulation, we compare model results domain-average model results from the parent domain against radio-soundings from the Praha-Libuš station - Please note, the meteorological station is positioned located roughly 11 km south-southeast from of our area of interestand

- 580 . Praha-Libuš is in an area with slightly different topography and urban topology located on at the southern edge of the city, which means that the comparison with model simulation is not exact and comparison with the model simulation cannot be exact and, especially within the lower parts of the boundary layer, modelled and observed profiles need not necessarily matchdue to the different surroundingscannot be expected to match. To estimate the spatial variability of the atmosphere between these two locations and thus to assess whether the soundings can be reliably used for evaluation of the PALM profiles, the WRF modelled
- 585 profiles in both locations, the sounding location and the PALM area, are provided.

Figure ??-5 shows vertical profiles of potential temperature and wind speed from PALM together with the soundings for the 20–21 July (episode summer e2ease). Taking into account the limitations of this comparison, for temperature the model simulations show good agreement with observations, capturing the overall shape of the profile with a slight tendency of underestimation of the actual values in the summer episodeto underestimate actual values. However, in the lower layers the

590 model tends to underestimate the diurnal variations, showing lower stability during the night and lower instability during the day.

Vertical profile of potential temperature (solid line) and wind speed (dashed line) for the summer e2 episode from the weather balloon soundings at Praha-Libuš station (purple and red) and parent domain average from the PALM model simulation (yellow and blue).

- 595 The wind speed generally follows the driving WRF profile except near the surface, where the wind speed tends to exhibit lower values due to increased surface friction from the explicit representation of microscale terrain features, buildings, and tall vegetation. During the first night (Fig. **??**5) the modelled and observed temperature profiles agree wellnear the surface, while further above the modelled profiles show slightly smaller values, although the shape of the profiles is similar to the observed one. The modelled wind speed in the residual layer is much smaller than in the measured profile and remains
- 600 approximately constant until noon while the measured wind slows down and becomes closer to the simulation during the daygenerally lower than the radiosonde. On the following day, the modelled and observed potential temperature profiles agree fairly wellup to 1,500 myery well, both indicating a vertically well mixed boundary layer. At that point in time the modelled profile indicates a higher boundary-layer top at about 1,950 m, while the observed profile indicates a boundary-layer top at about 1,550 m. However, as already mentioned above, this discrepancies can also be due to the different location
- 605 rather than a misrepresentation of the model, where the boundary-layer depth over the more strongly heated built-up area may be deeper than over suburban areas (Brugger et al., 2018). During the second night, the modelled profile indicates a cooler boundary layer which is less stable near the surface. On 21 July at 00:00, the wind speed profile agrees well with the measurementmeasurements. However, at 06:00 the low-level jet is still present in the observations but missing in the simulation. On the following day, again the modelled and the observed temperature profiles agree, although the modelled boundary
- 610 layer tends to be cooler by about 1 K. The wind speed is almost constant and uniformly smaller weakens during the day and is lower than the observations throughout the entire depth of the model domain.

Figure **??** <u>6</u> shows the modelled and observed profiles of potential temperature and wind speed during the for 4–5 December (episode winter e3episode). During the first night the temperature profile suggests a more pronounced stable boundary layer. On the following day the modelled temperature profile agrees fairly well with the observed profile. However, the shape of the wind

- 615 speed profile differs from the measured one. Notably, the boundary-layer values differ considerably. On the second night and the following during the second day the temperature profiles agree reasonably well, even though the modelled profile indicates a slightly warmer near-surface layer by of about 1 Kcompared to the observed profile. Considering the entire period, wind speed mostly matches the WRF-modelled profiles above 200 m but with some notable discrepancies compared to observations. Near the surface, PALM shows lower wind speeds compared both to the observations and WRF. At this point, however, we
- 620 would like to emphasise again that a direct comparison between the PALM-modelled profiles and the observation should be taken with care, especially within the near-surface layer where the profiles can be significantly affected by the different local surroundings.

Vertical profile of potential temperature (solid line) and wind speed (dashed line) for the winter e3 episode from the weather balloon soundings at Praha-Libuš station (purple and red) and parent domain average from the PALM model simulation (yellow 625 and blue).

3.6.1 Spatial development of the urban boundary layer

Figure ?? presents the time series of the wind speed and wind direction at the roof of the highest building of the inner LESdomain (FSv - Faculty of Civil Engineering CTU). The average wind speed in PALM agrees with the observations.

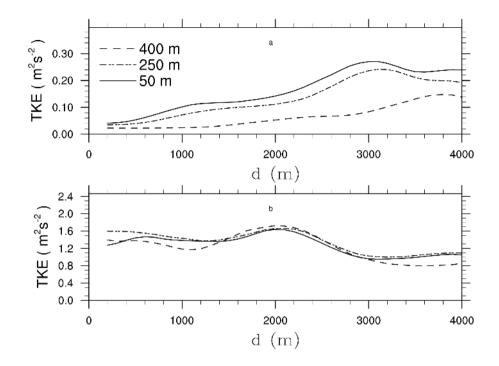


Figure 8. Time series of wind speed and wind direction at the roof of the tall building Horizontal profiles of 30-min time-averaged resolved-scale turbulent kinetic energy (TKE) in the Faculty of Civil Engineering of parent domain plotted against distance from the Czech Technical University inflow boundary (d) for summer episode e2 and a) the winter episode e3. Thin lines represent 10-minute averages case at 14 UTC on 5 December and b) the thick lines 1-hour moving averages of wind speedsummer case at 13 UTC on 20 July. The arrows represent 2-hour averages of wind directionTKE is shown for heights at 50 m, 250 m and 400 m above the terrain surface.

In the summer e2 episode the agreement is very good. The disagreement in the vertical profile at midnight of 20 July might be explained by the difference of timing of As described in Sect. 3.1, the parent domain receives boundary conditions from WRF where turbulent structures are not explicitly resolved. To trigger the spatial development of turbulence in the LES, synthetic turbulence is imposed at the lateral boundaries (Kadasch et al., 2020). However, even though this accelerates the development of turbulence in the LES, it still requires sufficiently large fetch distances for the turbulence to be spatially fully developed. Lee et al. (2018) have pointed out that an insufficiently developed turbulent flow can bias results in urban boundary-layer simulations. Hence, in order to assess how the turbulent flow develops within the model domain, Fig. 8 shows horizontal profiles of the turbulent kinetic energy (TKE) in the parent domain as distance from the inflow boundary increases. The TKE was computed as $TKE = 0.5 \cdot \sum w'_iw'_i$, with $w'_iw'_i = \overline{w_iw_i} - \overline{w_iw_i}$; the overbar denotes a 30-min temporal average. For each grid point we determined its distance to the inflow boundary for the given wind direction. In doing this, we calculated backward trajectories from the mean wind direction and determined the distance between the sampling location and the intersection point

- 640 of the backward trajectory with the closest inflow boundary. Further, variances were averaged over similar distances to the inflow boundary; we then sorted similar distances into equally-sized bins of 100 m to obtain a sufficiently large sample size for each discrete distance. Furthermore, we note that the TKE is evaluated at relative heights above the surface. In the winter case, which is characterized by neutrally-stratified conditions at the given time point (see Fig. 6), the TKE increases with increasing distances from the inflow boundary at all illustrated heights and peaks at about d = 3000 m in the sharp wind speed decrease
- 645 as suggested by surface layer, while the peak position at larger heights is shifted towards larger distances. In the summer case, which is characterized by convective conditions at the given time point, the TKE is approximately constant up to 2 km from the inflow boundary and then slightly decreases with further increasing distances. However, the heterogeneous orography and nature of the buildings means local effects that will also play a role so we would not expect to obtain a constant equilibrium TKE value. Taking into account that the child domain inflow boundary is placed at about 2 km from the parent inflow boundary
- 650 in both cases, turbulence has already been developed at the child domain boundary, so we are confident that the error made by the too short adjustment fetch length is minor, though we emphasize that especially for the winter case larger horizontal extents of the FSv time series. Also in the morning of 21 July when the vertical profiles show significant disagreement in higher levels the difference at FSv is much smaller. There is one extraneous sharp peak of wind speed in the afternoon of 21 July in the PALM results. 22 July is simulated very well including the morning and evening secondary maxima. The wind direction 655 changes on 19 and 22 July are captured well with some disagreement in the timing.

In the winter e3 episode the difference is larger. The morning of 4 December shows disagreement in the wind direction followed by a sharp extraneous peak in wind speed. After a well-predicted decrease in the morning of 5 December the wind speed somewhat overpredicted and there is another large extraneous peak in parent domain are also desirable in order to better represent mixing processes in the upper parts of the evening of 6 Decemberboundary layer. Moreover, the turbulent flow

- 660 depends on the upstream surface conditions (terrain, buildings, land-use, etc.) which in turn depend on the wind direction. With insufficiently large model domains such effects might not be well represented. However, the wind direction does not agree. This confirms the disagreement of the wind profiles in Fig. ??. The as our validation study mainly focuses on the building layer where turbulence is produced by building-induced shear, we believe the error induced by not completely representative upstream conditions is small and does not significantly affect our validation results.
- Beside the transition of the turbulent flow in the parent domain, the flow also undergoes a transition after entering the child domain with its finer grid resolution as discussed in detail in Hellsten et al. (2020). In order to evaluate whether turbulence has been sufficiently adapted within the child domain at locations where simulation results are compared against observations, Fig. 9 shows frequency spectra of the TKE at different distances to the inflow boundary. We sampled time series of the wind speed and wind direction for episodes summer e1, winter e1, and e2 are presented in supplements in-velocity components
- 670 at different positions over one hour and calculated the spectra for each sampling location; afterwards we averaged over all spectra with similar distance to the inflow boundary. In the winter case the spectra close to the inflow boundary show a significant drop-off of energy at smaller frequencies compared to spectra at distances ≥ 250 m, indicating that especially the smaller scales are still not sufficiently resolved on the numerical grid, while at larger distances no dependence on the sampling location can be observed. In the summer case the flow transition from the coarse into the fine grid is even faster; even

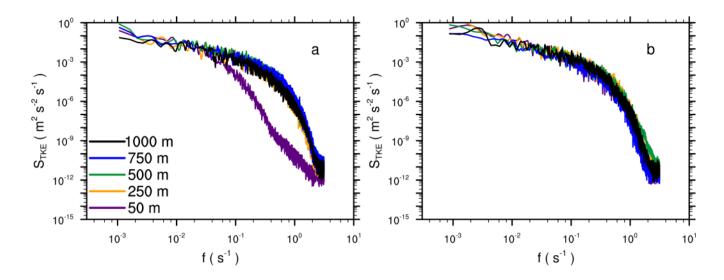


Figure 9. Frequency spectra of the TKE within the child domain at z = 50 m above the surface evaluated at locations with different distances downstream of the inflow boundary for a) the winter case at 14 UTC on 5 December, and b) the summer case at 13 UTC on 20 July.

675 spectra close to the inflow boundary indicate similar turbulence properties compared to the locations farther downstream. This is also in agreement with the findings presented in Hellsten et al. (2020) that under convective conditions the transition is small compared to neutrally-stratified or stable conditions as TKE is mainly produced locally by buoyancy rather than by shear.

3.7 Air quality

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For the CAMx model evaluation, urban background air quality monitoring stations closest to the PALM parent domain were
 used (see Sect. 2.3.6). Validation was performed for hourly average concentrations of NO_x and PM₁₀. Evaluation was done for all PALM simulation episodes which were then grouped as summer and winter. Metrics according to Britter and Schatzmann (2007) and Chang and Hanna (2004) for both campaigns are summarized in Table 1. For graphs of diurnal variation plotted by openair package (Carslaw and Ropkins, 2012) see Fig. \$17\$18.

For NO_X, the metrics show a significant underprediction of the measured concentrations (FB ca. -0.8) for both summer and
winter episodes. Nevertheless, the diurnal variation is captured quite well, although in winter modelled peaks in the evening are larger than in the morning while the reverse is seen in the observed data.

Summer PM_{10} concentrations are less underestimated with FB ca. -0.5, and morning and evening peaks are sharper and appear about 1 h earlier than in observations. Winter PM_{10} are even slightly overestimated but the CAMx model is not able to represent their real diurnal variation. Modelled diurnal variation is very similar to that for NO_X, which indicates that it is dominated by diurnal variation of traffic, while in reality different sources play a important role as well.

Table 1. Evaluation of CAMx 1-h concentrations against urban background stations for summer and winter episodes.

	NO	X	PM ₁₀			
	Summer	Winter	Summer	Winter		
Ň	684	816	<u>907</u>	1078		
<u>mean obs (μg.m⁻³)</u>	22.6	59.5	22.1	30.4		
$\underbrace{\text{mean mod } (\mu g.m^{-3})}_{\text{mean mod }}$	10.1	24.4	13.4	33.3		
FB	-0.76	-0.84	-0.49	0.09		
NMSE	1.51	2.15	0.65	0.53		
FAC2	0.38	0.31	0.50	0.69		
R	0.54	0.28	0.34	0.13		

N = ensemble size; mean obs = observed mean value; mean mod = modelled mean value; FB = fractional bias; NMSE = normalized mean square error; R = Pearson correlation coefficient.

4 Results

4.1 Surface temperature

In the following section we will discuss the model performance with respect to the surface temperature. First, we will show general surface temperature results and show an example of direct comparison against observed values. Then we will draw a broader picture of model performance for different types of surfaces, supported by relevant statistical measures. Subsequently, particular cases at individual locations will be presented and the related shortcomings of the model and the observations, as well as the implications of the shortcomings of the fine-scale input data, are discussed.

4.1.1 Overall performance

Figure 10 shows an example of a 3D view of instantaneous surface temperature in the child domain on 20 July 2018 at 14at 13:00 CETUTC 20 July. The heterogeneous distribution of surface temperature reflects the distribution of pavements and green

areas, with higher temperatures over paved areas and at building walls and roofs. Below the trees, where most of the shortwave direct radiation is absorbed within tree crowns, surface temperatures of about 290 K are modelled (e.g. on the right side of the figure or within courtyards), while higher surface temperatures up to 330 K are modelled at intensively irradiated vertical buildings building walls. Moreover, the effect of different wall and roof material parameters on surface temperature can be

705 identified, e.g. at with roofs showing lower surface temperatures where green fractions are present, whereas some while some other walls and roofs show values up to 320 K. In order to evaluate the modelled surface temperature more quantitatively, we compare the modelled surface temperature against observed values in the following parts of this section.

The observations cover a wide range of the street canyon configurations and surface types. A complete set of comparison graphs of the surface temperature for all points in all observation locations (see Fig. 1 in Sect. 2.1) for the summer e2 episode

710 of the observation campaign (19–21 July 2018) and for the winter e3 episode (4–6 December 2018) is given in the supplements



Figure 10. Example 3D view of the child modelling domain with 2m resolution from west-south direction. The colour scale represents the modelled surface temperature south-west direction on 20 July 2018 at 1413:00 CET_UTC (1314:00 UTCCET). The colour scale represents the modelled surface temperature.

in Sect. S3. As a supporting information, the graphs of the modelled values of surface heat flux (shf), ground heat flux (ghf), net radiation (rad_net), and incoming and outgoing shortwave (SW) and longwave (LW) radiation (rad_sw_in, rad_sw_out, rad_lw_in, rad_lw_out) are also available in the supplements in Sect. S4. The following subsections demonstrate the behaviour of typical urban environments and selected typical objects of the urban canopy.

715 4.1.2 Typical urban environments

The first selected typical situation represents a narrow street canyon of width between 10 and 20 m surrounded by traditional block-of-flats buildings with 4 to 6 stories over 20 m high (loc. 07, 08, 09; combination of LCZ 2 and 5 according to Stewart and Oke, 2012). Fig. **??** shows the situation for the location 07-2. The observations in the approximately south-north oriented street were done in the direction towards the west oriented (slightly south-inclined) wall.

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Observation location 07-2: the view of the observation location and IR and RGB photos with placement of the evaluation points (upper), observed (dots) and modelled (lines) surface temperature for wall (left) and ground (right) evaluation points for summer e2 (middle) and winter e3 (bottom) episodes. Top left image © 2020 Google.

In general, the modelled daily cycle of surface temperature agrees well with the observed surface temperature during the summer and winter episode. However, some discrepancies can be observed, e.g. during the summer e2 episode where the

725

modelled surface temperature at 07-2_H (pavement) shows slightly larger amplitude of the diurnal cycle (colder during the night, warmer in the daily maxima). At vertical surfaces, the modelled surface temperature agrees fairly well with the observed

one during the summer episode. Throughout most of the winter e2 episode, the modelled temperatures at both horizontal and vertical surfaces are generally overestimated (up to 5 K on walls), where the model does not capture the night cooling on the first night and the model remains systematically above the observations for the rest of the episode. The horizontal surfaces also show a sharp peak around noon on 4-Figure 11 shows an example for the observed and 5-December, which does not appear on

730 show a sharp 6 December.

The next typical situation represents a wide street canyon with a combination of the traditional and the contemporary buildings (loc. 01, 02, 06, 11). Fig. **??** shows surface temperatures sampled at loc. modelled diurnal cycle of surface temperature profiles at one particular evaluation location, 11-1, along with a street view of the location area and the RGB and IR views

- 735 of the location with the evaluation points labelled. 11-1 which is located on Evropská třída , the street, a west-east oriented boulevard of width between 40 and 50 mm in width (building to buildingwidth). The evaluation points are-), with evaluation points placed on the concrete tramway belt, pavement, and on the nearly south oriented wall of two traditional five-floor brick buildingswith, the left of which has an additional thermal insulation layerin the case of the left one. For the summer scenario, the modelled surface temperature agrees fairly well at the horizontal and vertical locations with respect to the daily-diurnal
- amplitude and temporal evolution. However, at the horizontal surfaces the modelled nighttime surface temperatures are underestimated by about 3-4-3-4 K. When the sun comes-up comes up the next day the modelled and observed surface temperature agrees fairly well again , meaning that surface temperature again matches the observed surface temperature, so the nighttime bias in surface temperature does not propagate to into the next day <u>simulation</u>. In the winter case the modelled surface temperatures also agree with the observations. At the second night where the surface temperature continuously decreases, the model
- 745 shows values higher by 1–2 K, except for the nights where the modelled surface temperatures are about 1–2 K higher than the observed ones at both horizontal and vertical surfaces, while the modelled surface temperatures at horizontal surfaces match quite well with the observations in the following day and the third night. At the vertical locations, however, . Further, two sharp peaks in the modelled daytime surface temperatures show two sharp peaks during the morning hours as well as during early afternoon which were not captured by the observation. It strikes that the modelled surface temperatures at the vertical
- 750 surfaces show also larger values (by about 2 K) compared to the observation during the following night. However, compared to the summer case, the deviation hours are striking, and not present in the observations. Similar peaks can be also observed at some other locations mainly during the winter episode. For a detailed discussion concerning these peaks we refer to Sect. 4.1.3 where this effect and its causes are analysed.

A complete set of modelled and observed diurnal cycles of surface temperature for all evaluation points in all observation locations (see Fig. 1 in Sect. 2.1) for the summer e2 episode (19–21 July 2018) and for the winter e3 episode (4–6 December 2018) is given in the supplements in Sect. S3. As supporting information, the graphs of the modelled surface temperatures to the observed one is in a similar range values of the surface sensible heat flux, ground heat flux, net radiation, and incoming and outgoing shortwave and longwave radiation are also available in the supplements in Sect. S4.

Location 12 covers a courtyard environment, representing another typical urban environment. The respective courtyard has
 a horizontal extent of approximately 160 m × 70 m and a depth of more than 20 m with large plant canopy inside the courtyard cavity. The observation points , depicted in Fig. ??, are placed on the south-southwest oriented wall of the Sinkule house

student dormitory, which was built before World War II and has been insulated in the meantime. The ground points are located on the asphalt/concrete playground near the building, the point 2 is influenced by large nearby trees. The model shows good agreement with the observations at the vertical surfaces. The observations cover a wide range of surface types. Since we cannot

- 765 show daily cycles for all observation points we condensed the results to show the general performance of the ground and wall modelling capability of PALM. To distinguish model behaviour for different types of surfaces, the evaluation points were put into the following categories: pavements (paved areas without traffic), streets (paved areas with traffic), grass, wall of traditional building, wall of contemporary office building, wall of building with glass or glass-like surface, and plant canopy affected surface. The complete assignment of the evaluation points to the particular categories is given in table Table S7. Figure 12
- 770 shows scatter plots of the modelled and observed surface temperature for particular surface types during the summerepisode, except for a slight underestimation of the diurnal cycle. Also at the horizontal surfaces the modelled surface temperature agrees well with the observed one, however, at nighttime the model underestimates surface temperature again. Compared to the summer case, the modelled surface temperature agrees less with the observedone for the winter episode. Here, the peak values at noon as well as the nighttimevalues are overestimated similarly to the previously shown two locations (see Fig. ??) and Fig. ??), which correlates with the overestimated global net radiation in this episode.

Location 12-1: the view of the observation location and IR and RGB photos with placement of the evaluation points (upper), observed (dots) and modelled (lines) surface temperature for wall (left) and ground (right) evaluation points for summer e2 (middle) and winter e3 (bottom) episodes. Top left image © 2020 Mapy.ez.

- The last typical group of the urban environment which occurs in the studied area is represented by the relatively open places, such as a square or a park (locations 03, 04, 05, 10). Fig. **??** shows examples of the three types of pavement and grass on the central square Vítězné náměstí. The surface temperature at e2 episode. The best agreement can be observed for street and pavement surfaces, and traditional building walls. At lower temperatures (which corresponds to nighttime values) the scatter is generally lower compared to higher surfaces temperatures, where, especially at the pavement surface in location 10-1 is captured very well by the model during the summer episode, while during the winter episode the model tends to
- 785 be warmer for the first two days of the observed period with overestimation of both daily minima and maxima. During the third day (6 December)the model is very close to the observations. The location 10-3 represents an example of buildings, a wide scatter can be observed. To support this qualitative impression from the scatter plots, Table 2 provides statistical error measures. Modelled surface temperatures at pavements and streets are slightly too cool, especially at nighttime, as indicated by the negative bias. Further, the RMSE indicates higher uncertainty at daytime and lower uncertainty at nighttime, especially
- 790 at building walls. The main reason for this behaviour is probably the typically lower thermal conductivity in comparison with ground surfaces, which causes more rapid reactions of the surface temperature to the changes in radiative forcing. This effect, in connection with binary changes in direct radiation during the course of the day due to shading effects, along with possible geometrical imperfections in the discretized terrain and building model, can cause temporally and spatially limited strong discrepancies between modelled and observed point values. This issue is analysed in more detail using the example
- 795 of location 11-1 V in Sect. 4.1.5. Mismatch of shading can also be caused by the imprecise description of the natural-like lawn. The diurnal cycle shapes of the tree crowns (see Sect. 4.1.6). Modelled surface temperatures at grass-like surfaces

also show good agreement with the observations, with mostly low scatter both during the day and at night, but with slightly overestimated nighttime values. A wider scatter, even at lower temperatures, can be observed for both glass-like surfaces and contemporary buildings walls, with the largest RMSE in the daytime. The reason for this higher spread is probably a more

- 800 complex wall structure and the higher uncertainty in its identification (see Sect. 4.1.3). In the case of glass-like surfaces these causes are accompanied by the fact that the IR camera photos of such locations contain a substantial amount of reflections from other surfaces (opposite buildings, sky), and so does not provide an adequate measure of the surface temperatureagrees with observations well in both summer and winter episodes with slight underestimation of daily maximum on 19 July and overestimation of night temperatures during both summer nights.
- 805 Location 10-1 and 10-3: the view of the observation locations and IR and RGB photos with placement of the evaluation points (upper), observed (dots) and modelled (lines) surface temperature for summer e2 (middle) and winter e3 (bottom) episodes.

4.1.2 Grounds modelling

The following sections describe behaviour of particular selected urban layer surfaces and objects across locations. These effects 810 are discussed in detail in Sect. 4.1.4.

Pavements and streetsbelong to the most frequent types of ground surfaces in the urban environment. For a complete list of corresponding evaluation points see categories *pavements* and *streets* Similarly, Fig. 13 shows scatter plots for episode winter e3. Again, at streets, pavements, grass-like, and traditional wall surfaces the scatter is relatively low though it does not show large difference between daytime and nighttime (see also RMSE in Table S7. The agreement of model results with

- 815 observations in case of pavements and streets is generally very good in most of the evaluation points. As an example see e. g. locations 10-1_H (Fig. ??), 11-1_H (Fig. ??), and point 12-1_H1 (2), in contrast to the summer case. In general, it is striking that in the winter case modelled surface temperatures are slightly overestimated, as indicated by the positive bias values. This is especially true for glass-like materials which show far too high modelled surface temperatures as well as a large scatter. However, the problems of surface temperature measurements of glass-like surfaces by IR cameras due to direct reflection
- 820 from other surfaces, which is mentioned above and discussed in detail in Sect. 4.1.4, applies here. Grass surfaces modelled temperatures are also overestimated. This overestimation can be seen in many individual locations (see supplements Sect. S3). The reason for this overestimation of surface temperatures, which is more pronounced in wintertime (compare Fig. ??). Notable exceptions12) than in summertime, however, are some locations influenced by trees (e.g. loc. 08, see remains unknown at this point. There is further discussion of modelling grass surfaces in summertime and the necessary pre-requisites below in

825 Sect. ??)4.1.2.

Modelling of grass-covered surfaces represents a challenge.

4.1.2 Grass surfaces

The energy balance of a grass-covered area may strongly depend on soil-water content, assumed plant cover, LAI, etc., which leaf-area index (LAI), along with other factors (Gehrke et al., 2020), and these are mostly unknown in this study. A complete

Table 2. Statistics of observed and modelled surface temperatures (K) for the simulation episodes summer e2 and winter e3.

		All times			Daytime			Nighttime		
Surface type		MB	MAB	RMSE	MB	MAB	RMSE	MB	MAB	RMSE
Pavements	S _c	<u>-0.7</u>	2.1	2.7	-0.2	2.3	3.0	-1.6	1.7	2.0
	w	1.5	1.7	1.9	1.9	2.0	2.4	1.4	1.5	1.7
Streets	S _~	-1.6	2.5	3.2	-1.4	2.7	3.6	- <u>2.1</u>	2.1	2.3
	W	0.9	1.0	1.4	1.3	1.4	2.0	<u>0.7</u>	$\underset{\sim}{\overset{0.8}{\sim}}$	1.0
Grass	S _c	0.6	2.7	4.1	0.3	3.2	<u>4.9</u>	1.1	1.7	2.0
	W	1.2	1.5	2.1	1.5	1.9	2.7	1.0	1.3	1.8
Walls (trad. building)	S_	-0.5	2.0	3.3	-0.3	2.5	3.9	-0.9	1.1	1.4
	W	1.7	1.9	2.6	2.2	2.3	3.5	1.5	1.7	2.0
Walls (contemp. building)	S.	-0.1	5.5	<u>7.4</u>	-0.4	<u>6.4</u>	<u>8.8</u>	0.2	4.2	4.5
	W	<u>4.9</u>	5.1	<u>6.8</u>	5.8	<u>6.3</u>	<u>9.6</u>	<u>4.5</u>	4.5	5.1
Walls (glass-like)	S_	1.9	3.6	5.3	1.8	4.2	<u>6.2</u>	<u>2.1</u>	2.6	3.2
	w	7.1	7.1	<u>7.9</u>	<u>6.8</u>	<u>6.8</u>	<u>7.8</u>	7.2	7.2	<u>8.0</u>
Plant canopy affected	S_	-0.8	2.5	3.6	-0.7	2.8	<u>4.1</u>	-1.0	1.6	1.8
	w	1.0	1.5	1.9	1.3	1.7	2.1	0.9	1.4	1.7

S = summer e2 episode; W = winter e3 episode; MB = mean bias; MAB = mean absolute bias; RMSE = root mean square error.

- 830 list of evaluation points placed on grass surfaces is provided in Table S7 in the category grass. Here, we examine points Let us examine three grass covered points, evaluation point 3 (EP 3) at location 05-1_H3, H, EP 2 at location 06-3_H2, and H, and EP 1 at location 10-3 H1 during the H during the second day of the summer e2 episode(Fig. ??) which, 20 July (see Fig. 14 and Sect. S3 of the supplements for detailed information on these particular locations). These points are not significantly influenced by any adjacent tree or wall . They and thus they are not affected by possible imperfection of the radiative transfer in the model. These points represent examples of three different types of grass conditions grass-type surfaces. The first of them, 835 $05-1_H3$ -point is placed in a recently built park with an integrated irrigation system, the second one $06-3_H2$ lays is located in a green tram line with a shallow soil layer and without any watering, and the third point $\frac{10-3}{10-3}$ H1 is located in a quite large lawn in the an open square area with a deep soil layer without watering and thus, and so resembling natural grass conditions. The daily cycle of the modelled and observed surface temperature at the different grass surfaces agrees fairly well with maximum temperatures of 35 °C, 52 °C, and 45 °C, respectively. Initial conditions of the soil were adopted from the WRF simulation 840 which represents spatially aggregated values over various surface types. To model To account for local differences in soil conditions properly for summer simulations, the grass areas within the model domain were split into three categories: naturallike grass, watered grass, and an urban grass type, while and the original WRF soil moisture was roughly adjusted by factors of 1.0(e.g. 10-3 H1), 2.0(e.g. 05-1 H3), and 0.5 (e.g. 06-3 H2), respectively for summer simulations respectively. Since we have
- 845 no information about soil moisture at that level of detail, the chosen adjustment factors are a best guess based on a survey of

the locations and personal experience. The soil moisture for winter simulations was not adjusted. Fig. The diurnal cycle of the modelled and observed surface temperature for the different grass surfaces agrees fairly well with maximum temperatures of 35 ?? also shows daily °C, 52 °C, and 45 °C, respectively. Figure 14 also shows diurnal cycles of surface temperature at point 06-3 H2 (tram line) and 05-1 H3 (irrigated park) these points from a test simulation where the soil moisture of grass surfaces

- 850 was uniformly prescribed from the WRF simulation. With non-adjusted soil moisture, the daytime surface temperature at point urban grass (location 06-3_H2 and, EP 2) and watered grass (location 05-1_H3, EP 3) is under- and overestimated compared to the observed oneobservations, respectively, while though it agrees fairly well for the adjusted soil-moisture case. This indicates that using correct soil moisture values is a necessary prerequisite to sufficiently model natural-type surfaces within adequately model grass-like surfaces within an urban environment. Sensitivity of the model For additional details concerning the sensitivity
- 855 of surface temperatures modelled by PALM to the initial soil moisture is investigated in in urban environments we also refer to Belda et al. (2020). Apart from soil moisture, sensitivity of grass surface temperatures on to other parameters such as LAI, plant cover, root-distribution, etc., might also be importantand needs further investigation. For details in this regard we refer to Gehrke et al. (2020) who studied the sensitivity of the energy-balance components to different soil as well as land-surface parameters.

860 4.1.3 Wall surface modellingComplex structure of the walls

Buildings represent an essential element of the urban canopy and they influence the street canyon environment substantially. Heterogeneity of building types and properties is usually higher than in case of ground surfaces . Heterogeneity of properties of building walls also very high due to windows, ledges, balconies, stuccoes, and other facade objects. Moreover, determination of building wall and roof properties is usually a very difficult process even in case of a detailed field survey. Nevertheless, we can roughly distinguish a few basic building categories in the studied domain.

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The first category represents traditional buildings built from insulated bricks or construction blocks or poured concrete with various types of plaster and occasionally provided with additional surface insulation. A complete list of evaluation points placed on such types of walls is given in Table S7 under the *wall* (traditional building) category. The agreement of model results with observations for this type of walls is usually very good and occasional deviations can be attributed to the inaccuracy of the

- 870 wall parameters. As an example, In the case of vertical surfaces ("walls"), the model behaves well for most cases of walls of traditional buildings, while walls of contemporary office buildings are modelled less accurately (see Fig. ?? shows the surface temperature for evaluation point 06-4_V. For the summer episode the modelled and observed surface temperature agree fairly well. For the winter episode the situation is different. On the first day the modelled and observed surface temperature agree.During the first night, however, the wall cools down less rapidly in the model compared to the observation, being about
- 875 5 °C warmer in the model, which is in accordance to the other locations where WRF / PALM is not able to capture the nighttime cooling. Even though the modelled surface temperature is overestimated on the following day and night, the daily cycle and amplitude of surface temperature is well captured by the model on the following day, indicating that PALM is able to sufficiently simulate the physics at such wall surfaces.

Location 6-4_V: IR and RGB photos of the observation location with placement of the evaluation points (left), observed (dots) and modelled (lines) surface temperature for summer e2 (top right) and winter e3 (bottom right) episodes.

- 880 (dots) and modelled (lines) surface temperature for summer e2 (top right) and winter e3 (bottom right) episodes. The next group of buildings includes the modern type of large buildings characterized by steel or steel/concrete skeleton and prefabricated walls with a complex layer structure. A complete list of evaluation points of this type is provide in Table S7 in category *wall (contemporary office building)*. The agreement of model results with observation is variable and usually lower for this type of surfaces. This can be attributed to a 12 and Fig. 13). We are convinced that the reason for this is the more
- 885 complex structure of wall layers which is not these walls which can not be fully described by the four layers allowed by the current version of the BSM wall heat model. AdditionallyPALM input standard. Moreover, gathering precise information about this type of structure proved to be more difficult. An example evaluation for points 02-3_V1 and quite difficult. Let us show an example using evaluation points 2 and 3 at location 02-3_V2 is shown in V (see Fig. ??.While the point 02-3_V1-15 and Sect. S3 of the supplements for full information about the location).
- 890 While point 1 is captured by the model quite well except for slight overestimation during the night and morning hours, the point 02-3_V2_point 2 evinces an overestimation of around 15°C during the afternoon hours. A closer direct inspection of this wall revealed that the wall it consists of a thin upper outer layer followed by a 10 -cm thin air layer and then followed by cm layer of air before the rest of the wall structure, while the model considered this in the model all this is considered as a continuous wall. The observed outermost layer thus was was thus cooled from both sides, an effect which was not modelled properly by the current captured by the wall model.

Location 02-3_V: IR and RGB photos of the observation location with placement of the evaluation points (left), observed (dots) and modelled (lines) surface temperature for summer e2 (top right) and winter e3 (bottom right) episodes.

4.1.4 Glass surfaces

- A third important category of the buildings in the modelling domain is represented by contemporary office buildings with 900 similar skeleton as in previous category but with surface constituted of glass or glass-like materials. Evaluation points belonging to this type are summarized in Table S7 as the category *wall (glass like surface building)*. As an example, the location 11-2_V is shown in Fig. ??. At daytime the modelled surface temperature agrees well with the observed one, whereas at nighttime the model shows significantly higher surfacetemperatures compared to the observation. However, we note that such a kind of building presents. Some buildings have walls covered with glass or similar types of reflecting surface. These walls present a
- 905 challenge for both observation and modelling. The surfaces of the main problem is the fact that the surfaces of these buildings are more or less specularand this fact causes that the IR camera observations contain, which means that a substantial part of the LW radiation reflected from an opposite object as can be seen on IR image in entering the IR camera is a reflection of whatever is behind the camera. For example, location 11-2_V (see Fig. ??.For example, the glass surface 16 and Sect. S3 of the supplements for full information about the location) is a north-facing building, the lower part of which has a glass surface. The
- 910 area of the building around the evaluation point 11-2_V2 evaluation point 2 reflects the sky , while around the evaluation point 11-2_V3 the glass surface reflects the opposite building (into the camera, while the area around evaluation point 3, located just below, reflects the building opposite into the camera (the building opposite is around location 11-1_V). Consequently, the

derived values of the surface temperature represents primarily the surface temperature of the reflected object (wall, ground, treetop, sky), not of the observed object itselfand thus.. This can be well demonstrated by the different observed values at points

915 EP 2 and EP 3. The modelling of this type of building cannot be considered as validated by this studythus cannot be validated by means of IR camera temperature measurements.

From the modelling point of view, specular reflections are not considered by the The modelling of the surroundings of these points can be partly influenced by the fact that the current version of RTM , which considers all reflections to be considers all surfaces as Lambertian (see Krč et al. 2020). This affects means that they reflect radiation in all directions in the model while

- 920 in reality, part of the radiation undergoes specular reflection according the law of reflection. This fact does not directly affect the reflecting surface itself but it can influence the distribution of reflected SW and LW radiation among nearby surfaces. As the amount of incoming direct radiation is significantly larger than the incoming reflected radiation (direct radiation can reach up to 900 W m⁻² while the reflected radiation is limited to 200 W m⁻² for most common cases), this effect has usually little practical impact and is masked by the effect of the direct radiation. An example of this effect can be seen in point 06-1_H2
- 925 in comparison with point_location 06-1_H3-H by comparing evaluation point 2 with EP 3 (see Fig. ??17 and alternatively supplements Sect S3 for full information about the location). These points are placed on similar asphalt concrete surfaces but with different distance to the nearby glass facade. While the surface temperature at the more-distant point 06-3_H3-3 is modelled well, it is underestimated point 2 is overestimated by about 7 °C at point 06-3_H2 on 20 July 2018 at hours 11-13 UTC, which we attribute to the specular reflection of . The observation at point 2 at these times shows an atypical increase
- 930 of about 7°C which is not observed at other points placed on the same surface type. We can attribute this lift to the effect of the specular reflection from the glass facadewhich. As this effect is not considered in the model. In reality, point 06-3 by the model, the model gives similar results for both points EP 2 and EP 3. Results for EP 1 (limestone pavement) are less affected by the missing specular radiation in the model due to its much higher albedo.

4.1.5 Rapid changes of surface temperature

- 935 Some of the graphs of the surface temperature contain strong "peaks" in the diurnal cycle of the modelled wall temperature (see e.g. Fig. 11). This effect can be seen mainly during the winter episode e.g. at locations 6-4_V, 7-1_V, 7-2_V, 8-2_V, 9-2_H2 receives more SW radiation by specular reflection of the direct sunlight from the glass facade before noon, so that the pavement can heat up.However, by assuming that all reflections are Lambertian in the RTM, we do not account for the additionally reflected SW radiation from the glass facade, resulting in less available energy to heat up the surface. The other point 06-1V6,
- 940 9-2_H1 (limestone pavement) is less affected by V7, 11-1_V). Similar peaks can be observed in the corresponding radiative, surface, and ground heat fluxes (see Sect. 4 in supplements). Some of these peaks can also be found in measurements (clearly visible e.g. for 6-4_V), though most observations contain no corresponding peaks. Let us analyse in more detail location 11-1_V (see Fig. 11) where this effect is very strong for evaluation points 1, 2, and 3 on 5 December.
- Figure 18 shows the observed IR and RGB camera photos at corresponding observation times together with modelled outerparts at the closest saved model timestep. For easier orientation, Fig. S19 in the missing specular radiation due to its much higher albedo, supplements shows an overview of the modelled surface temperatures in the given area at the same

time-steps. Fig. 19 provides the complete timeline of 10-minute model outputs of the wall surface temperature from 5:28 to 12:48 UTC. The time-steps shown in the previous figure Fig. 18 are highlighted with a red frame, the red dots denote the position of the evaluation points 1, 2, and 3.

- 950 The first peak takes place between the first and second observation times (7:51 and 9:26 UTC) and thus it does not appear in the IR observations. The situation of the second peak is more complicated. This peak partly overlaps with the fourth observation at 12:48 UTC, which is only reflected in the observations by a very small increase of the surface temperature at point 1. The reason for this can be seen in the comparison of the shading from direct radiation in the RGB photo and the corresponding figure for the modelled SW radiation (see Fig. 18). The shade created by the building on the opposite side of the street is
- 955 approximately 3 meters lower in the model than in reality at this time. These differences can be attributed to the geometrical imperfections of the digital building elevation model (BEM) used, as well as to the errors introduced by its discretization and by the PALM process of the placing of the buildings on the terrain. One of the sources of the imprecision in BEM can also be peripheral objects on the roof area (e.g. banisters, air-conditioning systems) which create shading but are not considered in BEM (see street view of shading buildings in Fig. S20 in supplements).
- 960 Validation of wall heat flux is discussed Figure 20 shows a detailed graph of location 11-1_V for times from 7:00 to 14:00 UTC and provides additional information about the diurnal cycle of the surface temperature at this location. The evaluation points 1, 2, and 3 correspond to points from the graph in Fig. 11 while new evaluation points 4, 5, and 6 were added on the top layers of the wall. The graph shows that the diurnal variability of the surface temperature in this location has similar magnitude in the model as in the observations. This supports our conclusion that the model (namely the radiative transfer and
- 965 surface energy balance) works reasonably well and the differences in the values at particular evaluation points and times can be attributed mainly to the geometrical imperfections of the model which produce differences of shading of the direct radiation. These changes of surface temperature also cause rapid changes of the temperature gradient in the wall which explains the peaks in the surface and ground heat flux visible in the corresponding graphs in Sect.4.2 and it provides additional information for the wall modelling- 5 of supplements. The positive and negative peaks in the ground heat flux correspond to start and end
- 970 times of irradiation of the given point by direct radiation. This analysis also suggests how complicated a problem is represented by spatially and temporally detailed modelling of radiation energy processes and the surface energy balance in the complex heterogeneous urban environment.

4.1.6 Plant canopy modellingeffects

Trees and shrubs are modelled in PALM as the resolved plant canopy (PC) which is described by a fully-3D structure of leaf area density (LAD). Beside affecting the turbulent flow by adding LAD-dependent drag, resolved plant canopy also affects the radiative transfer by partially intercepting SW and LW radiation as well as emitting LW radiation (see Krč et al. 2020). Further, the absorbed incoming radiation is transformed into latent and sensible heating terms which are considered within prognostic equations of potential temperature and humidity. Many evaluated points are more or less affected by some evaluation points are affected to different degrees by PC. A list of evaluation points where a significant impact of PC can be seen is given in

980 Table S7 in row plant the row "Plant canopy affected surface. Here". In this section, we focus only on the summer scenarios

since deciduous trees (which constitute the majority of the trees in the domain) carry no leaves during the winter. Impact of the branches during the winter episodes is roughly modelled as 10 % of the summer LAD.

The first two examples (see Fig. ?? and ??)show a situation of two-

Figure 21 shows two examples of locations affected by trees (locations 12-1_H and 08-2_H; for full information about

- 985 these locations see Sect S3 in supplements). Location 12-1_H is on the left, with two evaluation points placed on the same surface (cobblestone and asphalt concrete, respectively)where pointasphalt concrete). The direct radiation at point 2 is directly influenced by the tree shading while tree influence on the point influenced by tree shading but the tree shade does not reach evaluation point 1 is much lower this time of year. The shading of the treetop decreases the surface temperature after noon, which is well considered captured by the model.
- 990 Location 02-1_H: IR and RGB photos of the observation location with the placement of the evaluation points (left) and observed (dots) and modelled (lines) surface temperature for summer e2 episode (right).

Location 12-1_H: IR and RGB photos of the observation location with the placement of the evaluation points (left) and observed (dots) and modelled (lines) surface temperature for summer e2 episode (right).

The last example (Fig. ??) shows asphalt /concrete and cobblestone surfaces placed A similar situation at location 08-2_H is shown on the right. The evaluation points are similarly placed on an asphalt concrete surface in a street canyon with surrounded

- 995 shown on the right. The evaluation points are similarly placed on an asphalt concrete surface in a street canyon with surrounded by two alleys of trees with linked treetops forming an umbrella-like covering. The street surface temperature at location 08-2. H is underestimated by the model by up to 5 °C. While these surfaces are modelled well in Because a similar type of surface is modelled well at 12-1. H and other locations, a possible the most probably explanation for this discrepancy could be is the tree shading. The reason could be a general overestimation of LAD in the input data and/or a discrepancy in its spatial
- 1000 distribution. The large tree crowns tend to arrange themselves into clusters with free space between them (see e.g. Mottus, 2006). Figure 21, with spots of direct shortwave radiation as Fig. ?? passing through the canopy, and location views in Sect S3 of the supplements suggests that this is also the case . As this fact was not taken into account in generation of tree input data, the resulting homogeneous LAD overestimates the shading of the radiation in the simulation case at location 08-2_H. However, the method used for calculation of the LAD distribution within the tree crown does not consider such clusters, leading to possible
- 1005 underestimation of total transmissivity of the whole tree crown. Moreover, PALM uses a constant extinction coefficient for calculation of the optical density from the LAD value, which can lead to overestimation of optical density if clusters are significant at the subgrid scale. However, this can be mitigated by decreasing the LAD value. These examples confirm the importance of the precise estimate of the structure of the tree LAD in the inputs for the PALM simulations, though gathering of such information presents a complicated task.
- 1010 Location 08-2_H: IR and RGB photos of the observation location with the placement of the evaluation points (left) and observed (dots) and modelled (lines) surface temperature for summer e2 episode (right).

4.1.7 Discretization issues

PALM discretizes the domain in a Cartesian grid where all values in every grid box are represented by one value. This leads to standard discretization errors. Moreover, the current version of PALM uses the so-called mask method to represent obstacles

- 1015 (terrain, buildings), where a grid box is either 100% fluid or 100% obstacle and consequently any surface is represented by orthogonal grid faces (see. Fig. 4). Besides implications with respect to the near-surface flow dynamicsand an increase of effective roughness, a, which can be locally affected, this discretization increases effective roughness and enlarges surface area. The step-like representation also increases the total surface area, surface representation also modifies the direction of the normal vector and the mutual visibility of the particular grid surfaces, which in turn also affects the surface net radiation
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implications on for radiative transfer and surface energy balance.

The first observed consequence of the discretization is the fact that the subgrid-size surface features cannot be represented, while in reality, these objects can significantly influence the shading of the wall (e.g. Fig. ??). This effect parts of the surface. This effect can be observed in many of the studied locations (see Sect. S3 in supplements) and it needs to be carefully taken into account in when making point comparison of the related surface values.

and thus the surface energy balance. The observations of the surface temperature allow us to demonstrate a few selected

Camera RGB (left) and IR (right) photo of the wall of the building in the Zikova street (location 01-1_V) on 20 July 2018 at 10:13 CET.

The next important effect The effects caused by the step-like representation surface representation include artificial shading and the alteration of the surface is the artificial shading. This effect normal vector. Both these effects can be observed in case

- 1030 of "slope" terrain or walls ("and studied in the case of slope" in case of walls means terrain as well as in the case of non grid-aligned walls, i.e. walls which are oriented in one of the south-west, south-east, north-west, and north-east directions and which are approximated by step-like structures). Fig. **??** shows an example of the modelled surface temperature in the park around location 05-1_H. Terrain of the park is slightly declining in the south direction which causes three artificial "steps" in the gridded representation of the terrain, one near the building, next in the centre of the park, and last near the south side
- 1035 of the park. The induced artificial shading causes significant drop of the surface temperature in affected surface grid cells which is well visible in the figure and which has no counterpart in reality. Other grid surfaces which are not directly influenced by artificial shading are also affected by the discretization of the terrain. The gridded ground surface is oriented horizontally while the real terrain surface is inclined about 4° to the south. This inclination decreases the incoming direct SW radiation by about 4 % in case of situation presented in Fig. ?? and even more for other hours of the day when sun elevation angle is
- 1040 lower. As a consequence, the incoming radiation and surface temperature in the model are overestimated for these surfaces. A demonstration of this effect can be done on evaluation points 03-1_H1. As an example, 04-1_H1, and 05-1_H2. These points lay on exactly the same type of the surface and none of them is directly influenced by any building or tree. While the modelled and observed surface temperature agree well in the first two cases, the modelled temperature in point 05-1_H2 is overestimated about 4 °C during the day (see Fig. ??), which supports our hypothesis that the missing sloped-surface representation in the model equation in the prodel causes this bias.

The park beside the building of Technical library (location 05-1) on 20 July 2018 at 12:30 CET. RGB (upper left) and IR (upper middle) photo of the location, 3D view on modelled surface temperature in this location (bottom) and an illustrative aerial view of the location. Top right image © 2020 Mapy.ez.

¹⁰⁴⁵ model causes this bias.

Observed (dots) and modelled (lines) surface temperature at points 03-1_H1 (left), 04-1_H1 (middle) and 05-1_H2 (right) for the summer e2 episode.

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An example of the transformation of the buildings from GIS data to a grid structure and of the impact of this structure on the resulting surface temperature for Sinkule dormitory (location 12-1) is shown in Fig. S18. The effects of the artificial shading as well as the alteration of the surface normal vector can be illustrated on let us show the wall around the observation location 07-1 (see Fig. ??22; for complete location information see Sect S3 in the supplements). This wall is oriented to the east with a

- 1055 slight inclination to the north. The upper pictures show row of the figure shows the observed photo on 20 July 2018 at 109:37 CET and UTC and the 3D view of the modelled incoming SW radiation on this wall at the corresponding modelling time step. The bottom pictures show row shows the same situation approximately one hour later at 1110:38 CETUTC. In the first case, all the wall is irradiated by the direct sun radiation and the model result shows the direct solar radiation while the model results indicate artificial shading of some grid faces caused by the step-like representation of the wall. The second case one hour later
- 1060 shows the situation when the wall is shadowed in the reality but some corresponding modelled of the corresponding model grid faces are irradiated by direct sun solar radiation due to their slight turn to the east directionin comparison with the real wall. Next two Two further consequences of the orthogonally gridded model surfaces are an altered distribution of the reflected

radiation and artificial self-reflections owing to the step-like terrain and wall representation. The first effect is difficult to demonstrate in the observed data due to less direct attribution of the reflected radiation to the particular source surface and its

- 1065 partial masking individual source surfaces and due to the partial masking of reflected radiation by the stronger direct radiation. The second effect can be best tracked on the sloped roofs but they were not observed within this observation campaign. It can also be demonstrated on the walls demonstrated e.g. on the wall around the location 07-2_V on the 20 July 2018 at 11:37 CET (see Fig. ??23). The wall is not irradiated in reality by the direct sun direct solar radiation at this moment as can be seen from the RGB photo. The south facing model wall grid faces grids of the model wall ("steps") are illuminated by the direct radiation
- 1070 and the radiation reflected from them consequently then irradiates adjacent grid faces turned to the west (oriented close to the original wall direction), an effect which has no counterpart in reality.

These potential sources of problems <u>especially</u> need to be considered, <u>especially</u>, <u>due to its_due to their</u> local nature, for <u>when making</u> point-to-point comparisons of modelled and observed quantities. However, <u>in case of when</u> averaging over larger areas, <u>one may expect that</u> these artificial effects to partially mutually compensate due to the fact that the global amount of the

1075 incoming direct and diffuse unchanged amount of incoming global radiation, which represents the strongest radiation forcing, is similar on the original and discretized surface. This will make these artificial effects less important for practical model utilization than in case of the model point evaluation radiative forcing. However, the differences in reflections can still lead to significant changes in the global energy balance of the surface.

To estimate the impact of the discretization on the averaged simulation results of the wall, we ran two idealized simulations of a street canyon. The simulation domain had 2 m grid resolution and it contained one west-east oriented 30 m wide street canyon of height 20 m. The simulated day was 19 July 2018 (the first day of the episode summer e2). The radiation was simulated by the coupled RRTMG model and the meteorological conditions were set constant (west wind 1 m s⁻¹, potential temperature at surface 295 K). The simulation started at 3:00 am with a preceding 24 hours spin-up run and covered the 16 sunny hours of the day. The first simulation employed the standard grid with no rotation while the second one had the grid

- 1085 rotated by 45° , utilizing PALM's ability to set grid rotation. This means that the walls of the street canyon were precisely aligned with the grid in the first case while they were represented by steps-like structures in the second case due to the 45° angle they form with the grid. The averaged results of the surface temperature, shortwave irradiation and net radiation over the south facing wall are presented in the supplements in Fig. S22 and S23. The results shows that the differences can reach about 3 °C for surface temperature, over 100 W m⁻² for shortwave irradiance and about 80 W m⁻² in the case of net radiation. These
- 1090 effects cannot be simply neglected and further more focused research is needed. Some potential amends in ways to amend the model are discussed in Sect. 5.2.

4.2 Wall heat flux

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The observations of the wall heat flux (HF) in two locations (see Sect. 2.3.2) allow direct comparison with the wall heat flux simulated by the model. Moreover, the observations of the surface temperature from the sensor allow to validate both validation of the PALM model as well as and the observations obtained by the IR camera (see Sect. 2.3.2).

During the summer campaign, the HF observations took place in Sinkule house from 19 July to 3 August 2018 and in location and at the Zelená location from 3 to 7 August2018... This period overlaps only partly with the modelling episode summer e2. The graphs of heat flux and surface temperature are shown in Fig. ??24. The sharp rise in observed HF and temperature before 06:00 UTC is caused by the direct irradiation of the sensors by the Sun and the data between around 6

- 1100 and 8 UTC cannot therefore be taken as valid measurements. (Similar peaks are visible in the PALM outputs before sunset.) The sharp drop of HF on 20 July after 06:00 UTC was caused by the sensor becoming unglued, which was fixed at about 08:00 UTC. The modelled and observed wall heat flux on the ground floor shows a similar daily cycle with a diurnal cycle with similar amplitude, though the model slightly overestimates the observed values by about 5 to 10W.m⁻². W m⁻², while the corresponding modelled surface temperature agrees fairly well with the observations. The modelled wall heat flux on the first
- 1105 floor shows a pronounced daily diurnal cycle, while the observed wall heat flux shows only a weak daily diurnal cycle with a significantly smaller amplitude. The modelled surface temperature, however, shows a smaller amplitude with higher nighttime but lower daytime temperatures compared to the observation observations, which is in agreement to with the respective wall heat fluxes where the model increasingly partitions the available energy into the wall heat flux.

The winter HF observations at the Sinkule house cover the episode e3 from 4–6 December 2018 and the observations in 10 location at the Zelená fit to episode location fit with the e2 in days episode for 27–28 November 2018 (see Fig. ??25). Even though the modelled surface temperature at the Sinkule house for the ground floor observation is slightly overestimated on the second day overestimated by around 2 °C with respect to the observed one during day time, the modelled and observed wall heat fluxes agree fairly well during the shown period period shown, especially for the first and second day. In contrast, on the first floor the modelled wall heat flux (absolute value) and surface temperature are strongly overestimated, especially during the

1115 nights. The minimum of the modelled wall heat flux reaches goes down to -50W.m^{-2} . W m⁻² during the night from 5 December to 6 December while observations suggest values between -10 and -15W.m^{-2} . W m⁻². The situation in location at the Zelená location is similar; the modelled observed HF fluctuates around -40W.m^{-2} during the nights while the modelled

counterpart reaches down to -80 W.m^{-2} goes down to $-80 \text{ W} \text{ m}^{-2}$. This behaviour suggests that the thermal wall resistance in the case of higher floors of the buildings Sinkule house Sinkule and Zelená buildings are underestimated. Both locations are older

1120 buildings Sinkule house is an older building which had been insulated in the past except for the ground floor. The real thermal resistance of this additional insulation layer, which is set in the input data to approximately 6 cm of polystyrene, is probably underestimated and the real insulation is more efficient. The details of the material of wall of Zelená building were not available and some type of construction block was assumed but its thermal conductivity in the model is probably overestimated.

4.3 Street canyon meteorological quantities

- 1125 Data collected by the mobile meteorological stations and vehicles allow us to compare the main variables characterizing the atmospheric properties inside modelled atmospheric quantities against observations within several street canyons. A large spatial variability inside a street canyon is expected and predicted by LES, therefore perfect agreement of the simulation and measurement cannot be expected. The comparisons also This section presents graphs and statistics of modelled and observed temperature and wind speed. Graphs are presented for summer e1, summer e2, and winter e3 episodes here; the complete
- 1130 results for all episodes are available in supplements Sect. S5 which also contains corresponding graphs of vertical sensible heat flux and relative humidity. The comparison graphs contain values from the WRF simulation to allow to assess the contribution assessment of the benefits of the micro-scale model. This section presents a comparison of the modelled temperature and wind speed with observations. Vertical sensible heat flux, relative humidity, and wind components can be found in the supplements in Fig. S19, Fig. S20, Fig. S21, and Fig. S22 respectively.

1135 4.3.1 Air temperature

Graphs of the air temperature at 3.9 m (Sinkule house) and 4.6 m (other locations) are plotted in Fig. 26Figure 26 shows timeseries of modelled and observed air temperature within different street-canyons for the summer e1, summer e2, and winter e3 episodes. In the summer campaign, the diurnal cycle is generally well resolved with daily maximum temperatures agreeing better than the daily minimum temperatures. The nighttime cooling is most often underestimated and hence the minimum temperature is too high. However, on certain days, e.g. on the 16 July, the simulated minimum temperature is lower than the observed one. The maximum temperature on the previous day was also higher than the simulated one. The comparison with the scenarios, the daily cycle of air temperature is generally captured by PALM. The modelled maximum air temperature generally agrees well with the observed maximum but is somewhat underestimated, especially at the Sinkule location. The

- modelled nighttime minimum values tend to be too warm compared to the observation, which is in accordance with the less
 stable modelled conditions as indicated by Fig. 5. The spatial variability of the shown model air temperature, as indicated by
 the red-shaded area, is rather low, suggesting that the comparison of modelled and observed air temperature do not suffer from
 any location biases. In addition, Fig. 26 also shows one-hourly averaged 2 m air temperature as modelled by WRF and inferred
 from the WRF-grid point closest to the observations. These values allow us to estimate whether deviations of PALM-modelled
 values arise from the driving synoptic simulation or from a different source. As WRF was set up without urban parameterization
- 1150 and no buildings were directly considered in the WRF 2m temperature in the closest point shows that on several days the WRF

modelled temperature agrees to the night street-canyon observations more closely. During the summer campaign the global radiation is well predicted by WRF except 23 July, which, however, coincides with the break in the episodes and movement of the measuring vehiclessimulation, a comparison between PALM and WRF with respect to street canyon temperature would thus not be expedient. Similar to the temperature simulated in PALM, the WRF-modelled 2 m air temperature also shows

1155 lower daytime maximum temperatures compared to the observations, while at nighttime even lower minimum temperatures are modelled which is in contrast to PALM. This, in turn, suggests that the too warm nighttime temperatures within the street canyon do not arise from the driving mesoscale simulation but for a different reason.

During the winter campaign, the simulated air temperature follows the observations less reliably. The behaviour changes during the simulated period. In episode 1, the daily peak temperatures are overpredicted. On the second day of the episode,

- 1160 the observed In the winter case, the modelled air temperature reflects the evolution of the observed air temperature, though the air temperature during the first day and the minimum temperature during the day is stationary while it significantly rises in the model. The cooling at first night is overestimated in all street canyons while WRF-modelled temperatures agree well with the observations. Starting from the second night until the end of the episode is captured correctly. However, it strikes that PALM mostly follows the daily cycle simulated by WRF.
- 1165 During the winter e2 episode, the temperature evolution is characterized by significant simulation, it is striking that the modelled air temperature is significantly overestimated by about 2 to 5 K. This can be attributed to the driving mesoscale WRF simulation which indicates a similar overestimation of air temperature when WRF was not able to capture nighttime cooling. This is reflected in the model, though with a much weaker intensity. On the first day the daytime temperatures are also overpredicted which is in accordance to the strong over-prediction of the global radiation by WRF (Fig. ??). PALM closely
- 1170 follows the temperatures in WRF during the episode. That can also be observed in episode c3 nicely shows that the performance of the building-resolving LES strongly depends on the driving mesoscale simulation. If the results on the mesoscale are biased this error will also propagate into the LES.

Statistical metrics for the model performance over all locations and scenarios considered are given in Table 3. For the summer scenarios the street-canyon air temperature is slightly underestimated during daytime, while it is overestimated during nighttime

- 1175 due to insufficient cooling near the surface. For the winter scenarios PALM overestimates the day- and night-time temperatures by about 1.5 K, which can be partly explained by the driving synoptic conditions. The scatter between observations and model results is about 2 K without any significant difference between day- and night-time or as summer- and winter-time. It is striking that the correlation between modelled and observed air temperature is higher during the daytime where the cooling on the night from 4 to 5 December is well reproduced except the last few hours. After well predicted daytime temperatures and an
- 1180 over-predicted night, the episode ends with an increase of temperature for 24 hours, which is overpredicted. That is again likely related to the over-predicted global radiation in WRF on 6 December (Fig. ??) daily cycle is usually well captured, whereas the correlation is less at nighttime where the nighttime air temperature is often overestimated.

Table 3. Statistical metrics of modelled one-hour averaged air temperature within different street canyons. The statistics are evaluated over all locations and episodes considered and are partitioned into summer and winter, as well as day- and night-time. The statistical metrics for the modelled 2 m air temperature in WRF are also given for completeness.

		Summer	episodes		Winter episodes				
	Day		Night		Day		Night		
	PALM	WRF	PALM	WRF	PALM	WRF	PALM	WRF	
\sim	233	233	122	122	210	210	370	363	
mean obs (°C)	24.1	24.1	<u>19.3</u>	<u>19.3</u>	3.5	3.5	2.4	2.4	
mean mod (°C)	23.5	22.4	<u>20.0</u>	17.7	5.1	4.2	4.0	2.7	
$\underbrace{\mathbf{MB}}_{\mathbf{K}}(\overset{\circ}{\mathbf{C}})$	<u>-0.6</u>	-1.7	0.7	-1.6	1.6	0.7	1.6	0.3	
RMSE (°C)	2.0	2.4	1.8	2.3	2.1	1.7	2.5	2.2	
R	0.91	<u>0.93</u>	0.73	0.78	0.91	0.89	0.85	0.81	

N = ensemble size; **mean obs** = observed mean value; **mean mod** = modelled mean value; **MB** = mean bias; **RMSE** = root mean square error; **R** = Pearson correlation coefficient.

4.3.2 Wind speed

The simulated and observed wind speed in the respective street canyons for the summercampaign generally show good

- 1185 agreement, even though they also indicate significantly larger wind speeds in the model at both Orlík locations (see episodes summer e1, summer e2, and winter e3 is summarized in Table 4 and plotted in Fig. 27). These locations also show large spatial gradients in the form of a large spread of the simulated wind speed in the neighbouring grid points (the shaded band in Fig. 27). That means that the spatial representativeness of the point measurement is limited and that the simulated values are very sensitive to the exact position of the sampling. The observing vehicle was located close to the wall. The large trees in the
- 1190 Terronská street, where ; the complete graphs for all episodes are shown in the supplements in Sect. S5. The graphs also show values simulated by the WRF model to illustrate the added value of the high-resolution LES simulations. Summary metrics for both models and all episodes (Table 4) show similar model performance in summer and winter with only slightly better statistics in summer. Both campaigns exhibit a significant over-estimation. However, all measures show that PALM is partially able to correct biases imposed by its driving boundary conditions.
- 1195 The wind speed in the summer campaign generally shows good agreement except at the Orlík stations were located, also increase the uncertainty, where the trees at top location, where significantly larger wind speeds are simulated by the model. We hypothesize that this is attributable to the nearby tree crowns in the street, which have a radius of 2 m in the model, but a radius of about 5 m in reality (see corresponding photo in Fig. S7). This discrepancy could have also influenced other modelled variables. For other locations, in the supplements). The uncertainty of the results in this location is also increased by large
- 1200 spatial gradients of the wind speed in the street and courtyard locations generally agree well with the observationsnear the buildings which makes precise fitting of the modelled and observed values sensitive to any spatial inaccuracy.

Table 4. Statistical metrics of modelled one-hourly averaged wind velocities within different street canyons. The statistics are evaluated over all locations and episodes considered. Summer and winter episodes are distinguished. The statistical metrics for the modelled 10 m wind speed in WRF are also given for completeness.

	Summer episodes		Winter e	pisodes	All episodes		
	PALM	WRF	PALM	WRF	PALM	WRF	
Ň	354	354	<u>580</u>	573	<u>934</u>	<u>927</u>	
$\underbrace{\text{mean obs } (m s^{-1})}_{}$	0.5	0.5	0.6	0.6	0.5	0.5	
$\underbrace{\text{mean mod }(m s^{-1})}_{\text{mean mod }(m s^{-1})}$	0.9	2.0	1.1	3.5	1.0	<u>2.9</u>	
FB	0.5	1.2	0.6	1.4	0.6	1.4	
NMSE	1.0	3.4	1.3	<u>6.6</u>	1.2	5.9	
R	0.50	0.38	0.55	0.45	0.53	0.42	

N = ensemble size; mean obs = observed mean value; mean mod = modelled mean value; FB = fractional bias; NMSE = normalized mean square error; R = Pearson correlation coefficient.

In the winter campaign the behaviour is more complicated. Episode 1 is very calm which is also reflected in the simulations. The comparisons at the Orlík stations in episode 2 are affected by the same factors connected with the spatial representativeness and tree size as in the summer campaign. Most of the episode 2 still shows a good agreement at Sinkule house, but the night to 29 November shows an increase of wind speedwhich cannot be observed in the measurements. The increase in wind speed is in accordance to the simulated wind speed in WRF and is connected with an occluded front passing the area. Episode 3 is

variable with moderate overestimations at certain times. The overestimations on 4 December can-

The daily cycle of the modelled wind speed in the winter scenario is roughly captured at the Sinkule location, except for the nighttime where the PALM-modelled wind speed is generally overestimated as also indicated by Fig. 6. This overestimation of

1210 the modelled wind speed, which is also accompanied by increased temporal variability, is also visible at the other stations; this might be linked to the insufficient representation of the stable boundary layer. Also the daytime values are mostly overestimated but this overestimation is much lower than that during nights. The overestimation in general could also be linked to the inaccuracies in the whole wind profile and hence the boundary conditions from WRF in which overestimates near-surface wind speed, which is expected when not using an urban parameterization, see e.g. Halenka et al., 2019, while also at higher levels the wind speed is partly overestimated (see Fig. ??6).

4.4 Street canyon air quality quantities

This section presents

1205

4.3.1 Wind speed on the roof

To assess model behaviour in the urban canopy outside the street canyon, a comparison of modelled and observed concentrations 1220 of NO_X . The results for PM₁₀ can be found the wind speed measured at the roof of the highest building in the child LES domain

Table 5. Comparison of 1-h average wind speed measured on the rooftop of FSv with WRF and PALM results for the same location.

	Summer episodes		Winter e	pisodes	All episodes		
	PALM	WRF	PALM	WRF	PALM	WRF	
Ň	176	172	219	213	<u>395</u>	385	
$\underbrace{\text{mean obs } (m s^{-1})}_{\text{mean obs } (m s^{-1})}$	2.3	2.3	1.7	1.7	2.0 ~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	2.0	
$\underbrace{\text{mean mod } (m s^{-1})}_{\text{mean mod } (m s^{-1})}$	2.5	3.5	2.6	4.1	2.5	3.8	
FB	0.07	0.41	0.43	0.85	0.26	0.65	
NMSE	0.34	<u>0.47</u>	0.75	1.47	0.54	$\underbrace{0.97}_{\ldots}$	
R	0.61	0.60	0.43	0.59	0.49	0.52	

N = ensemble size; mean obs = observed mean value; mean mod = modelled mean value; FB = fractional bias; NMSE = normalized mean square error; R = Pearson correlation coefficient.

(FSv - Faculty of Civil Engineering CTU) with PALM is presented. In order to illustrate the added value of the high-resolution LES simulations, outputs of the WRF are provided as well together with measurements from the nearest synoptic station Praha-Ruzyně for reference (reliable wind direction measurements were only available from the synoptic station). The graphs for summer episode e2 and winter e3 are in Figure 28. The time series for episodes summer e1, winter e1, and e2 are presented

- 1225 in the supplements in Fig. S23. Figure 29 shows the simulated and measured concentrations of NO_X S21. Summary metrics for all episodes are in Table 5. The wind speed is generally overestimated, with smaller errors in the summer and the winter campaign. In general, the modelled concentrations show a similar magnitude and indicate a similar temporal evolution as the measurements. It strikes that the simulated concentrations are simulations, a difference already present in the driving WRF simulation. In a comparison of the two models, PALM shows better agreement with observations with the exception of the
- 1230 correlation coefficient, which is similar in summer and even higher for WRF results in winter. For most of the episode, the PALM simulated wind speed is closer to the measurements in the summer cases, even though also in the summer cases concentrations temporally deviate from the measured values significantly, e.g. during nighttime and the morning hours on 15 July where concentrations are significantly underestimated. In the evening of 15 July, there was a large concentration peak simulated, but no peak can be observed in the measured data which becomes especially apparent at the Sinkule house
- 1235 station where the scatter of the observed concentrations is only small within the enclosed courtyard cavity. Another large overprediction of modelled concentration can be observed in the morning of 21 July. Also, for both these overpredictions, the CAMx mesoscale simulation shows larger concentrations compared to the measurements in FSv observations than the WRF results as well as the background Praha-Ruzyně observations. During the winter e3 episode, the courtyard. A possible reason for temporal mismatch of modelled and observed concentration might lie in the different wind speeds. The modelled
- 1240 wind speed profiles are significantly lower compared to differences are considerable. In particular, there is a large peak in the observed wind speed from the aerological soundings at evening of 6 December, which confirms the disagreement of the wind profiles in Fig. 6.

4.4 Street canyon air quality

This section presents a comparison of modelled and observed concentrations of NO_X and PM₁₀. The simulated and measured

- 1245 concentrations of NO_X in the summer e1, e2 and the winter e3 episodes are shown in Figure 29. The complete graphs for NO_X, PM₁₀, and PM_{2.5} for all episodes can be found in the supplements in Sect. S5. Summary statistics for NO_X 1 hour average concentrations for aggregated summer and winter episodes are presented in Tables 6 and 7. Statistics were calculated separately for street canyon locations influenced directly by the traffic and for the courtyard of Sinkule house, which, with respect to traffic, represents an urban background. Similar summary statistics for PM₁₀ are presented in the supplements (Table
- 1250 S8 and S9).

PALM coupled with a driving mesoscale model has a potential to represent both the magnitude and the temporal evolution of street level NO_X concentrations and thus eliminate the underprediction of the mesoscale model. This is especially true for different types of street canyons, but it is also important to mention that the differences between urban background and street canyon locations are captured well. Variability of PALM 1-hour average NO_X concentrations expressed as a standard deviation

- 1255 is about 50 % larger than that of observed data in summer episodes for both street canyon and background locations. In winter episodes the situation is opposite. When we check the large PALM ovepredictions (e.g. 15 July after sunset, 21 July in the morning, or 25 November after sunrise), these all happen, almost exclusively, when the driving CAMx model gives values within the range of, or even largely overestimates, the observations. Similarly, situations when PALM underpredicts NO_x concentrations happen when the increase in observed values is not reflected by the driving model as is the case for the 2nd half
- 1260 of the winter e2 episode. As can be seen from Fig. S17, a strong surface temperature inversion on 28 November at 00:00 UTC and especially 06:00UTC 21 July (Fig. 27), especially near the surface. At midnight on 21 July, both the modelled and the observed wind profiles indicate a well pronounced low level jet near the surface. Later during the morning hours, the low-level jet can still be observed in the measured profile, while it is not present any more in the modelled profile, accompanied with lower wind speeds and less mixing near the surface, UTC is not captured by the WRF which in turn favours the built-up of
- 1265 higher concentrations in the model impacts PALM meteorology (which at least partially reflects the observed inversion) and boundary concentrations.

Graphs of NO_X concentrations in street-canyon measuring locations. Yellow symbols denote observed 10-minute concentration averages, the black curve 10-minute concentration averages computed by PALM and the yellow curve 1-hour moving concentration averages computed by PALM. The light green band shows the interval between the smallest and the largest 10-minute average

1270 value among the neighbouring grid points. The red curve denotes the 1-hour concentration averages at the closest CAMx grid point.

In winter, the modelling of NO_X concentrations is complicated by It is also evident that the simulated NO_X concentrations are closer to the measurements in the summer episodes, especially in the street canyon locations. However, a high resolution modelling of concentrations in winter is more challenging due to local heating and the associated uncertainties of the emissions.

1275 The strong simulated peak in the morning of 25 November, which is also present in the CAMx results, does not appear to be

Table 6. Comparison of 1-h average NO_X concentrations measured in the street canyons with CAMx and PALM results for the same location.

	Summer episodes		Winter episodes		All episodes	
	PALM	CAMx	PALM	CAMx	PALM	CAMx
$\overset{\mathbf{N}}{\sim}$	224	<u>224</u>	363	<u>360</u>	<u>.587</u>	<u>584</u>
<u>mean obs (μgm⁻³)</u>	22.6	22.6	54.5	<u>54.7</u>	42.3	42.4
$\underbrace{\text{mean mod}}_{(\mu \text{gm}^{-3})}$	26.2	4.6	42.1	13.9	36.0	10.4
standard deviation obs $(\mu g m^{-3})$	14.9	14.9	<u>56.1</u>	<u>56.3</u>	47.7	47.8
<u>standard deviation mod ($\mu g m^{-3}$)</u>	21.8	4.5	33.4	13.8	30.5	.12.1
FB	0.1	- <u>1.3</u>	-0.3	-1.2	-0.2	- <u>1.2</u>
<u>NMSE</u>	0.5	5.0	0.8	5.5	0.8	<u>6.4</u>
FAC2	0.70	0.09	0.67	0.20	0.68	$\underbrace{0.16}$
R	0.62	0.29	0.70	0.52	0.70	0.57

N = ensemble size; **obs** = observed concentration; **mod** = modelled value; **FB** = fractional bias; **NMSE** = normalized mean square error; **FAC2** = fraction of predictions within a factor of two of the observations; **R** = Pearson correlation coefficient.

present in the measurements at all. A detailed examination of the concentration fields show a strong effect of the revealed a strong impact of local heating sources and also the effect of effects from the boundary conditions.

The validation metrics according to Britter and Schatzmann (2007) and Chang and Hanna (2004) for the summer and the winter campaign are summarized in Table ??. The statistics were For PM₁₀ PALM overpredicts observations during winter
1280 episodes and also the variability of its outputs is ca. 50 % larger than in observed data, the complete opposite of the case for NO_x.

<u>The PALM metrics for NO_X and PM₁₀ computed from all available 10-minute concentration averages in 1-hour concentration</u> averages at all points where measurements were available . The metrics (not shown) fulfil the criteria for dispersion models as suggested by Chang and Hanna (2004). Namely the Although these criteria were developed for simpler models, they are

- 1285 applied to a more complex problem here and are good indicators of fitness for purpose. More specifically, the absolute value of fractional bias is less than 30%, the modelled values fit within the interval given by the half and the double of the observed value then 0.3, the fraction of predictions within a factor of two of the observations is more than 50% of the time, and the random scatter is within the factor of one expressed as geometric variance (VG; not shown in tables) is within a factor of two of the mean value. The performance in the summer campaign and in the winter campaign only differs in the sign of the fractional
- 1290 bias which is negative in summer (overpredictions) and positive in winter (underpredictions) (i.e. VG < 1.6). These criteria are also fulfilled for data split into summer / winter episodes and street canyon / background locations with the following exceptions: VG is 1.8 for winter background NO_X and no criteria are fulfilled for summer background PM₁₀.

In addition to the stationary measurements, mobile observations of the air quality indicators were performed (see Sect. 2.3.4 for details). Here we compare Fig. 30 shows graphs comparing observed values of NO_X with modelled values in grid points boxes corresponding to the position of the mobile instruments for NO_X (Fig. 30; for . For comparison of PM₁₀ see Fig. S24 in the supplement). For the summer episode (19 July) the morning measurements are shown.

Table 7. The model performance evaluation metrics according to Britter and Schatzmann (2007) and Chang and Hanna (2004) computedfrom 10-minute average concentrations Comparison of 1-h average NOx modelled by concentrations measured in the Sinkule yard withCAMx and PALM results for the same location.

	summer S	Summer episodes	winter	Winter episodes	All ep	isodes
<u>N-</u>	3,039 - PALM	3,816 CAMx	PALM	CAMx	PALM	CAMx
mean obs g.m ^{-3} N	15.9 - <u>130</u>	4 8.3 _130	200	.197	330	327
$\frac{\text{mean mod g.m}^{-3}}{\text{mean obs }(\mu \text{gm}^{-3})}$	19.0-8.6	38.9 -8.6	33.9	34.2	23.9	24.0
FB -mean mod $(\mu g m^{-3})$	<u>-0.18-9.6</u>	0.21-5.7	35.5	13.6	25.3	10.5
NMSE standard deviation obs $(\mu g m^{-3})$	0.85-7<u>.1</u>	2.88- 7.1	<u>39.1</u>	<u>39.3</u>	33.1	33.2
FAC2 standard deviation mod $(\mu g m^{-3})$	0.69-9.5	0.64 5.5	29.5	12.7	<u>26.9</u>	.11.1
R−FB	0.59 0.1	0.63 - <u>-0.4</u>	0.0	-0.9	0.1	- <u>0.8</u>
<u>NMSE</u>	0.9	0.8	1.0	3.7	1.2	4.2
FAC2	0.78	0.60	0.66	0.49	0.71	0.53
R	0.50	0.62	0.54	0.39	0.61	<u>0.4</u> 7

N = ensemble size; **obs** = observed concentration; **mod** = modelled value; **FB** = fractional bias; **NMSE** = normalized mean square error; **FAC2** = fraction of predictions within a factor of two of the observations; **R** = Pearson correlation coefficient.

The observed Sect. S5 in the supplements. The observed NO_X values show quite high variability within the short timeframe of the measurement in many location measurements in many locations (variability between $20-160\mu$.g.m⁻³ µg m⁻³). On the other hand, the oscillations are very small during some other measurements (e.g. loc. 6–17 on 19 July or and partly loc. 13 on 4 December). This high variability of some measured values suggests impact of a very close local source of emission emission

1300 4 December). This high variability of some measured values suggests impact of a very close local source of emission emission source (e.g. bus on bus station buses at bus stations or local heating) which, however, but this cannot be verified with the data available. Moreover, these oscillations are not present in the PM_{10} observations, which supports the aforementioned hypothesis of local NO_x sources in contrast to dynamical causes.

In the winter episode, on 4 December NO_X observations show much higher variability than in the summer episode. During

1305 the morning series, modelled values correspond well with measurements quite well with observations for the most part with the exception of the measurement pointpoint 11 in which where the model is 2–5 times lower. The afternoon series shows good agreement in at points 2, 3, 4, 13 and 14. In At points 10 and 11 the model results are again consistently lower than in realityobservations.

5 Summary and conclusions

1310 5.1 Summary of the results

In this study, PALM LES simulations nested into the driven by mesoscale WRF and CAMx simulations were performed for a real urban environment in Prague-Dejvice, Czech Republic. Meteorological Modelled meteorological, air qualityand

wall-surface, and surface quantities were compared against in-situ measurements taken during a specially designed observation campaign. Air temperature, wind speed and chemical concentrations agreed well with the observations with respect to their

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The PALM model properly adjusts to the temporally evolving WRF and CAMx conditions and simulates the temporal evolution and daily amplitude , except for the observed strong nighttime cooling on 15 July, 20 July, and 21 July, which was not well captured by the LES, probably due to a misrepresentation of the stable conditions. This issue needs further investigation in the future. The modelled of street-canyon air temperature and wind speed agree well with the observations and properly adjusts

- 1320 to the temporally evolving WRF conditions. However, the modelled wind speed shows higher values compared to the observed one for location Orlik in some time periods, which can be explained by large spatial gradients near the buildings and by the tree crowns which are partly too small in quantities in most cases, with some noticeable exceptions such as insufficient nighttime cooling in some conditions. However, correct results depend on proper driving conditions as well as on the correct setting of the urban canopy properties in the model and the model. Further, especially during the winter episodes, meteorological quantities
- 1325 resemble the simulated WRF values due to a weaker local energy forcing, meaning that the accuracy of the model results is strongly related to WRFaccuracyspatial and temporal distribution of emissions. The importance of the urban canopy properties was demonstrated in many particular cases. It was also shown that the driving WRF model does not perfectly reproduce the observations, resulting in discrepancies in the wind speed and potential temperature propagating into the PALM solution via the boundary conditions given by WRF.
- 1330 Concentrations of NO_X were modelled well in some situations and PALM properly supplies the adds a local air pollution increment to the urban background values provided by CAMx simulation the CAMx simulation, while for some places and times (mainly about around sunset or sunrise) it overpredicts the model overestimates the concentrations of NO_X(e. g. 15 July evening, 25 November morning). That. This is probably related to atmospheric stability and uncertainties of in modelling stably stratified turbulent flow. The opposite situation (i.e. the underpredicting underestimation of NO_X) occurs less often (e.g.
- 1335 during the night from 14 July to 15 July). These discrepancies could be partially connected with attributed to uncertainties of the emission and with imperfection in emissions and imperfection in boundary conditions provided by WRF and CAMxprovided boundary conditions but a more, though another probable cause is the deviation of the PALM modelled PALM misrepresenting the turbulent flow under some meteorological conditions. This issue needs further investigation. PM₁₀ concentrations were modelled less accurately than NO_x, which can be mainly attributed to the driving model and overestimated emissions of
- 1340 <u>resuspended dust.</u>

The modelled surface temperature agrees <u>reasonably</u> well with the observed one at most of the surface <u>evaluation_evaluation</u> points. However, it <u>strikes is striking</u> that the agreement is usually better for the summer episodes when strong radiative forcing exists than for the winter episodes when the model results are more prone to uncertain <u>specification of</u> material properties as well as inaccuracies with respect to the atmospheric conditions given by the in atmospheric conditions from the driving

1345 mesoscale model. The surface temperature at pavement surfaces and at wall surfaces belonging to traditional buildings built e.g. from based on bricks or building blocks - is usually modelled well, while the surface temperature at modern buildings with multi-layer prefabricated walls is less well captured . At natural captured less accurately. At low-vegetated ground surfaces the modelled surface temperature agrees also well with the observation, even though we note that the model results strongly depend on a proper description of initial soil moistureand

- 1350 properly other surface-material, and probably other soil parameters. Beside an accurate prescription of surface-material parameters, also an accurate representation of the LAD is also essential for accurate modelling of the local atmosphere-surface exchange. Even though this study contains some indicative sensitivity investigations for the studied domain and episodes, we note that a systematic sensitivity study on the model input parameters is out of the scope of this paper and the reader is here referred to Belda et al. (2020). Furthermore, issues related to the discrete representation of the terrain and building surfaces on
- 1355 the Cartesian grid revealed to be a crucial factor for model inaccuracy (see Sect. 4.1.7).

5.2 Outlook of model development Lessons learned and data improvement outlook for future improvements

This study also points towards particular aspects in the model, its configuration, the input data preparationand, and the observation strategy that deserve particular focus in the future.

The current version of the PALM input standard (PIDS) and implementation of BSM allows discretization of the walls into four layers, independent of the thickness and the structure of the real wall, meaning that the grid resolution of the wall layers may differ among different wall surfaces. Further, wall material properties for complex walls with multiple layers are sometimes not well captured by only four wall layers, leading to under- or over-estimation of the thickness of the insulating layer, among other discrepancies. A variable number of wall layers would allow more realistic representation of wall material properties. Moreover, pre-specified typical structures of complex wall compositions in BSM would simplify proper initialization of these walls.

<u>The current method of discretization of terrain and buildings in PALM is bound to the Cartesian model grid, which means that an the entire volume of each grid cell contains either atmosphere (free or with plant canopy) or obstacle(terrain or building).</u> As a result, every model surface forms a boundary between grid cells and its normal is parallel to one of the grid axesor obstacle. If the modelled domain contains uneven terrain, sloped roofs, or walls that are not parallel to the grid axes, e.g.

- 1370 facades that are aligned along the NE-SW direction, the discretization creates artificial steps which affect radiative fluxes as well as the airflow. For example, such Such step-like surfaces create artificial shading or sunlit surfaces modifying on facades create both artificially shaded and artificially sunlit surfaces which also affects the energy balance on the microscale. Further, observation points cannot be assigned arbitrarily to the discrete grid any more, which in turn complicates the analysis for non-grid aligned facades. For such locations, the values of modelled variables may need further postprocessing, or may be
- 1375 even unsuitable for validation and a location further away from the step need to be used instead. Examples of these issues are presented in Section 4.1.7 of the facade. Even though these effects are strongest locally, they can also bias the aggregated values for larger surface areas. A major change of discretization is planned for future versions of the PALM model – surfaces will be represented using the Immersed Boundary Method (see Peskin, 1972). This method allows to represent representation of surfaces with arbitrary orientation, thus avoiding the negating creation of artificial steps.
- 1380 In the current version of the RTM model-radiative transfer model (RTM), all surfaces are considered as Lambertian reflectors, meaning that directional reflection at windows or polished materials cannot be considered, which, however, even though such

reflection can be found at almost every facade. This in turn adds uncertainty to the surface net radiation and thus to the energy balance at the surrounding surfaces. Implementation of specular reflection is planned to better simulate the radiative transfer at glass and polished surfaces.

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The current implementation of the BSM discretizes walls by four layers, independent of the thickness or the material of the wall, meaning that the grid resolution of the wall layers may differ among different wall surfaces. Further, wall material properties at walls with multiple layers are sometimes not well considered by only four wall layers, leading e.g. to an underor overestimation of the thickness of the insulating layer. A variable number of wall layers would allow to represent wall material properties more realistically. Moreover, pre-prepared typical structures of the complex wall composition in BEM would simplify proper initialization of these walls.

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The analysis of air and surface temperatures revealed insufficient cooling of the air during nights nocturnal air cooling in certain meteorological conditions where the stratification is underestimated not captured properly by the model. In this studywe explicitly prescribe, the incoming radiation where is explicitly prescribed, while radiative cooling of the air volume itself is not considered. Hence, in order to check how sensitive the model results are on this, we also ran to this, test simulations where we applied the RRTMG radiation scheme and where radiative cooling of the air volume is considered were run; however, we could observe observed a similar insufficient cooling in this case. This insufficient cooling during nighttime requires further investigation in the future nocturnal cooling requires further future investigation.

Another implication arises from the mesoscale nesting approach. The analysis of the wind speeds at higher levelsand temperatures, and of temperatures, revealed that PALM mostly partly reflects the conditions simulated by the mesoscale model (WRF), especially during wintertime. This in turn suggests that the model domain of the nested LESsimulation might have a small to develop its own equilibrium. However, this needs further investigation in the future. Measurements

- be too small to develop its own equilibrium. However, this needs further investigation in the future. Moreover, as the mesoscale simulation does not resolve the turbulent flow, we need to impose synthetic turbulence at the inflow boundaries. Even though the inflow is already turbulent, The error made on the mesoscale is thus propagated into the LES, biasing its simulation results. To minimize this mesoscale forcing bias on the LES results, the turbulent flow needs to develop spatially downstream of the inflow
- boundary forming coherent structures which require significant large fetch lengths of several kilometers (Muñoz-Esparza et al., 2017; Lee e . However, in the urban layer the adjustment is faster and 1-2 km are sufficient for the building-affected layer (Lee et al., 2018)
 Although the flow within the building-affected layer is well developed in the analysis area, the turbulent flow within the upper parts of the boundary layer has still not been fully developed. This implies that mixing processes at the boundary-layer top and its impact on near-surface microscale processes might be not considered well, though this was not the focus of this studydriving.
- 1410 mesoscale conditions might be further combined with additional nudging terms inferred from observations, continuously nudging the imposed boundary conditions for the LES towards the observations.

The study suggests strong sensitivity of the results on accurate to accuracy of input data, e.g. for such as the wall-material properties and the structure of tree crowns. The sensitivity of the PALM model to the PALM to material parameters is more systematically investigated in Belda et al. (2020). Bulk parameters prescribed for certain building categories might strongly

1415 deviate from the actual conditions at the building. Hence, usage of bulk input parameters might significantly modify the simulation results locally. Other specific detailed observations are needed to improve properties of the categories of wall, roofs,

and pavement materials. The study also stresses the need for precise boundary conditions as well as correct setting of the initial soil moisture for natural low vegetation surfaces.

- The experimental campaign also serves as a source of useful experience for future studies of similar type. Modern build-1420 ings with high amounts of glass and other reflective surfaces on the surface exterior surfaces proved to be challenging for surface temperature measurements using an IR camera. The reflections often hide the thermal radiation of the surface. A higher number of traditional buildings (bricks, concrete) would allow better assessment of the accuracy of the building surface parameterizations, which are primarily developed for these types of buildings and of the accuracy of the parameters assumed for these buildings. obscure the emitted thermal radiation from the surface and thus the IR camera does not provide a reliable
- 1425 way to observe surface temperature for such surfaces.

Data from mobile measurement vehicles proved to be difficult to interpret and difficult to draw statistically relevant conclusions from due to the influence of the strong local temporally and spatially evolving emissions, which are difficult to simulate in the emission model. In future, either a significantly higher number of measurements would be required or the effort should concentrate be concentrated elsewhere. One of the directions to be considered direction for consideration is a combination of

1430 traditionally full-featured comprehensive vehicle-observation stations with a network of the sensors. Further wider network of more limited sensors.

Though drones at first sight offer another promising direction, drone measurements in a city are <u>unfortunately</u> limited by various restrictions <u>based imposed</u> by the air traffic control and land owners. The <u>whole entire</u> city of Prague is located in controlled airspace that starts at the ground and starting at ground level and including our area of interestalso lies in restricted

- 1435 airspace. A drone flight must be programmed with regard to the properties of the measurement sensors, e.g. the relaxation time, and preparatory test flights. Other requirements for useful drone observations are matching height and speed changes to instrumentation characteristics, such as relaxation time. Preparatory test flights in consultation with the drone operator may be necessary. Regular aerological balloon soundings from the Praha-Libuš station proved to be indispensable. In future, increasing the frequency of measurements during a measurement campaign would be very useful and the possibility of dedicated
- 1440 soundings in the area of interest should be considered. However, this is also limited by restrictions similar to those on drone observations.

In summary, the ability of PALM to represent reality to a reasonable degree depends not just on the representation of physical processes in the model itself, but on input-data quality and the accuracy of the mesoscale forcing. For future studies it is thus a valid question where the focus should lie; should it be on further improving the model to better reflect physical processes in the

- 1445 urban boundary layer, or on obtaining as accurate and detailed input data as possible. In the authors' opinion, however, these options are not mutually exclusive but have to be balanced against each other. Focusing mainly on the input data will sooner or later result in a situation where the model performance is constrained by an insufficient representation of the physics, and a model with perfect physical processes will still need very good and detailed input data to produce practically relevant output. The task of attributing relative importance of these sources of uncertainty has been extensively tested in the field of numerical
- 1450 weather prediction and climate modelling in a number of coordinated projects producing large ensembles of simulations, e.g. the currently ongoing CMIP6 (Eyring et al., 2016) and CORDEX (Giorgi et al., 2009; Gutowski et al., 2016). In our case, a

similar approach of employing different models and model setups, and testing their respective sensitivity to input data would allow assessment of the sources of uncertainty. However, due to the enormous computational resources required for these kinds of simulations, such an endeavour is not feasible for one modelling team and it would benefit from the kind of framework of coordinated experiments that are a norm in the climate modelling community.

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Code and data availability. The PALM modeling system is freely available from http://palm-model.org (last access: 29 May 2020) and distributed under the GNU General Public Licence v3 (http://www.gnu.org/copyleft/gpl.html, last access: 29 May 2020). The model source code version 6.0 in revision r4508 used in this article is also available via https://doi.org/10.25835/0073713 (Resler et al. , 2020a). The configurations and inputs of the model for all simulated episodes are available via http://hdl.handle.net/11104/0315416 (Resler et al. , 2020b).

1460 Appendix A: Statistical measures used in manuscript

Apart from means and standard deviations of observed and modelled values, the following normalised statistics are used to summarise model performance. Please note that we adopted the convention that bias is positive when the model overestimates observations.

factor of two (FAC2): fraction of predictions within a factor of two of the observations

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fractional bias:

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$$FB_X = 2 * \frac{\overline{X_{\text{model}} - X_{\text{obs}}}}{\overline{X_{\text{model}} + X_{\text{obs}}}}$$

normalised mean square error:

$$NMSE_{X} = \frac{\overline{(X_{\text{model}} - X_{\text{obs}})^{2}}}{\overline{X_{\text{model}}} * \overline{X_{\text{obs}}}}$$

1470 For temperature given in degrees Celsius the following non-normalised statistics were used: mean bias:

$$MB = \overline{T_{\text{model}} - T_{\text{obs}}}$$

mean absolute bias:

$$MAB = \overline{|T_{\text{model}} - T_{\text{obs}}|}$$

1475 root mean square error:

$$RMSE = \sqrt{(T_{\text{model}} - T_{\text{obs}})^2}$$

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<sup>&</sup>lt;sup>1</sup>https://www.fona.de

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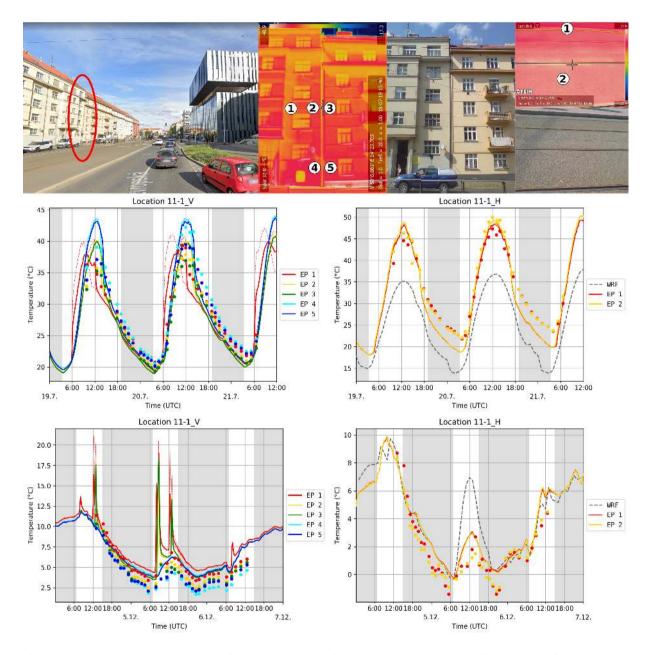
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**Figure 11.** Observation location 11-1: the view of the observation location and IR and RGB photos with placement of the evaluation points (upper row), and graphs of observed (dots) and modelled (lines) surface temperature for wall (left) and ground (right) for particular evaluation points (EP) for summer e2 (middle) and winter e3 (bottom) episodes. The modelled values come from the child PALM domain, the dotted and dashed lines represent the modelled temperature for left and right grid faces (see description at the beginning of this sectionSect. 4.1.1). The grey dashed line shows the corresponding WRF skin layer temperature for horizontal surfaces. The grey areas denote the night time. Top left image © 2020 Google.

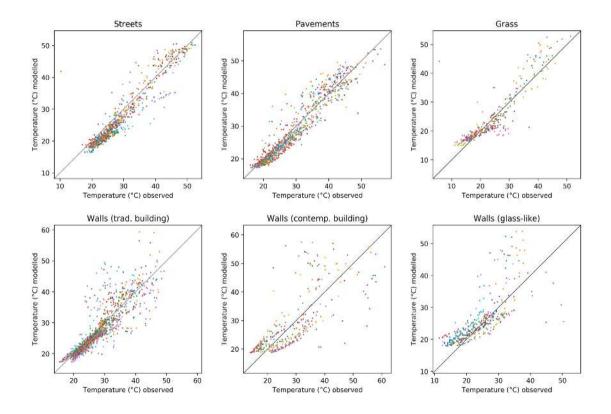
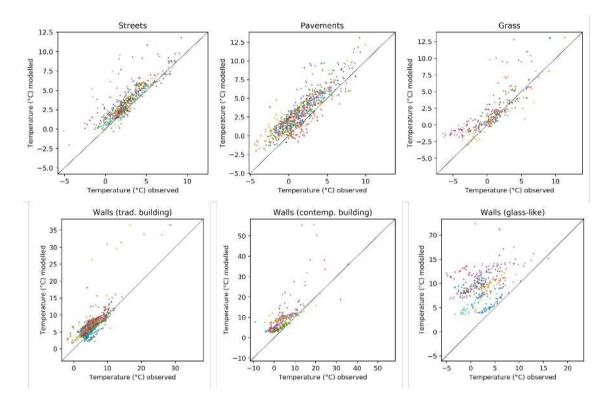
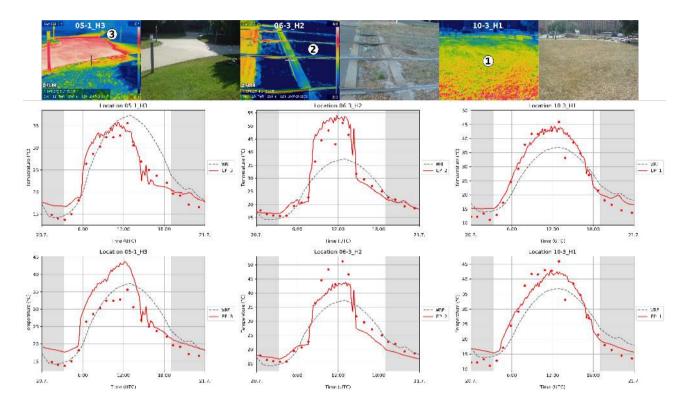


Figure 12. Scatter plots of the modelled and observed surface temperature for particular types of the surfaces in Table S7 during the summer e2 episode. Individual evaluation points are plotted by dots of different colours.



**Figure 13.** Scatter plots of the modelled and observed surface temperature for particular types of the surfaces in Table S7 during the winter e3 episode. Individual evaluation points are plotted by dots of different colours.



**Figure 14.** IR and RGB photos with evaluation points of locations 05-1\_H3H, 06-3\_H2H, and 10-3\_H1-H with placement of selected evaluation points (EP) (upper row) which represent three different grass types grass-type surfaces found in the modelled urban area. Observed The observed and modelled surface temperature at these locations . Middle right graph shows the time development of soil moisture in the first two PALM soil layers during the simulation for 20 July 2018 (second day of episode summer e2. Two lower graphs show , middle row), and the same results from testing a test one-day simulation with all grass surfaces initialized initialised with soil moisture uniformly prescribed from WRF simulation for evaluation points 05-1\_H3 and 06-3\_H2output (bottom row). The grey dashed line shows the corresponding WRF skin layer temperature. The grey areas denote the night time. All results are from the child 2m resolution domain.

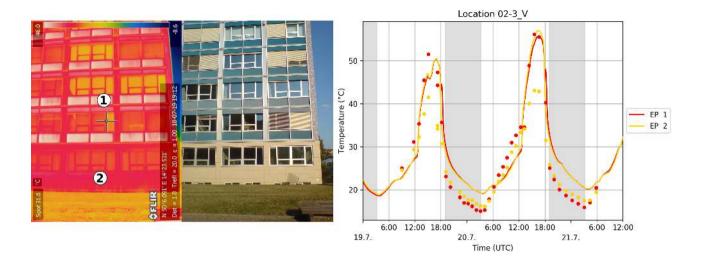
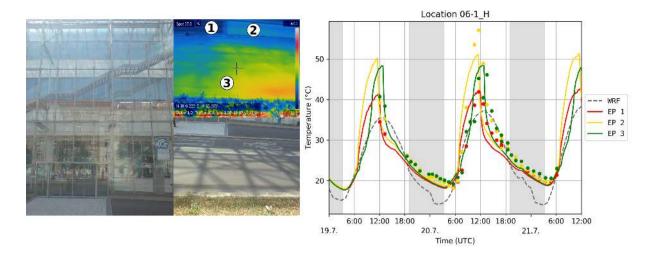


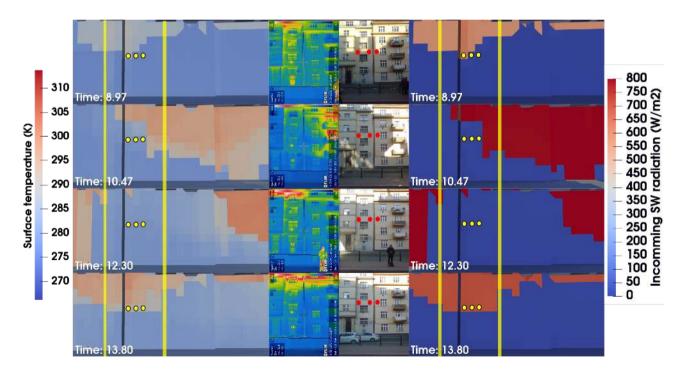
Figure 15. Observation location 2-3\_V: IR and RGB photos of the building with placement of the evaluation points (left) and graph of observed (dots) and modelled (lines) surface temperature for particular evaluation points (EP) for summer e2 episode (right). The grey areas denote the night time. Top left image © 2020 Google.



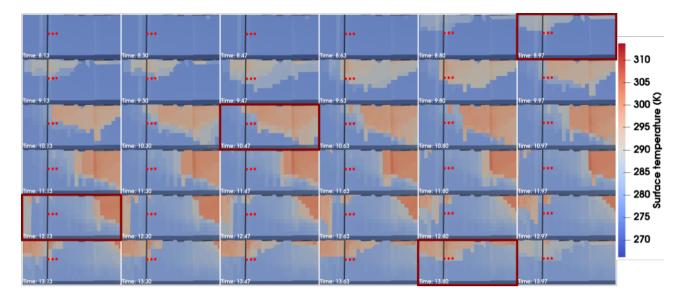
**Figure 16.** Location Observation location 11-2\_V: the view of the observation location (left), IR and RGB photos of the observation location building with placement of the evaluation points (leftcentre), and graph of observed (dots) and modelled (lines) surface temperature for summer e2 particular evaluation points (top rightEP) and winter e3 for summer e2 episode (bottom right)episodes. The grey areas denote the night time. Top left image © 2020 Google.



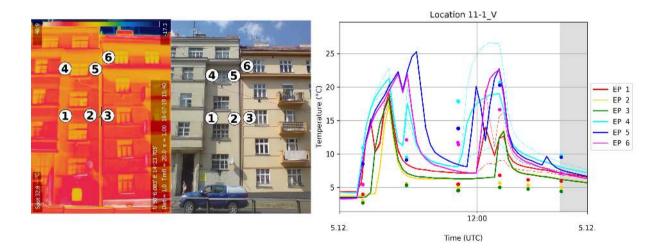
**Figure 17.** Location 06-1: IR and RGB photos of the observation location with placement of the evaluation points (top), observed (dots) and modelled (lines) surface temperature for summer e2 episode (bottom). The grey dashed line shows the corresponding WRF skin layer temperature. The grey areas denote the night time. Top left image © 2020 Google.



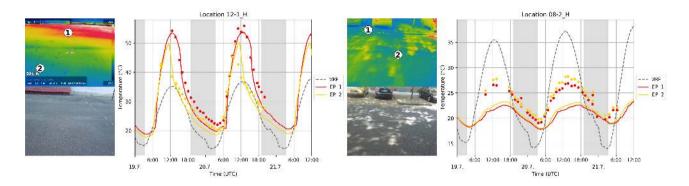
**Figure 18.** Observed camera photos (IR centre left and RGB centre right) on 5 December 2018 at observation times 7:51, 9:26, 11:18, and 12:48 UTC and the modelled counterparts for the closest saved model time-step: surface temperature (left) and incoming SW radiation (right). The yellow dots denote positions of evaluation points 1,2, and 3 (Fig. 11) and the yellow lines show the extent of the area shown on IR and RGB photos. For technical reasons, the step times for the model views express minutes as decimal fractions of the hours.



**Figure 19.** Timeline of 10-minute model outputs of wall surface temperature on 5 December 2018 from 5:28 to 12:48 UTC. The time-steps from the previous figure Fig. 18 are highlighted with a red frame, the positions of the evaluation points 1,2, and 3 are marked by red dots. For technical reasons, the step times for the model views express minutes as decimal fractions of the hours.



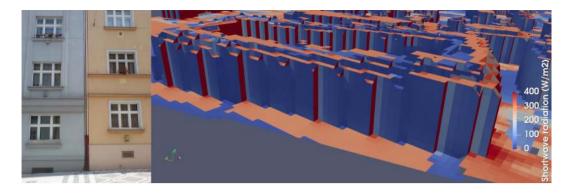
**Figure 20.** Location 06-1Comparison of IR observations and model at location 11-1\_V on 5 December 2018 from 7:00 UTC to 16:00 UTC. The left photos show IR and RGB photos images of the observation location with placement marked places of the evaluation points and the right image shows the graph of the modelled (topline), and observed (dots) and modelled (lines) values of the surface temperature for summer e2 episode (bottom) these evaluation points. Top left image © 2020 Google The grey area denotes the night time.



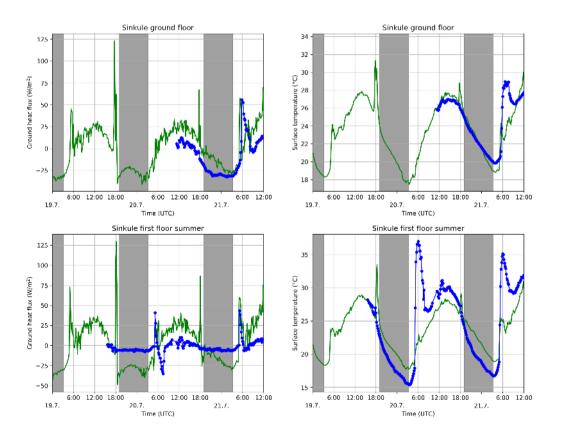
**Figure 21.** IR and RGB photos with locations of evaluation points and the graph of the observed (dots) and modelled (lines) surface temperature for these evaluation points during summer e2 episode (19-21 July 2018). The left half of the figure shows location 12-1\_H (the asphalt playground in the courtyard of Sinkule house) and the right part location 08-2\_H (asphalt concrete surface in Terronska street). The grey dashed line shows the corresponding WRF skin layer temperature. The grey areas denote the night time.



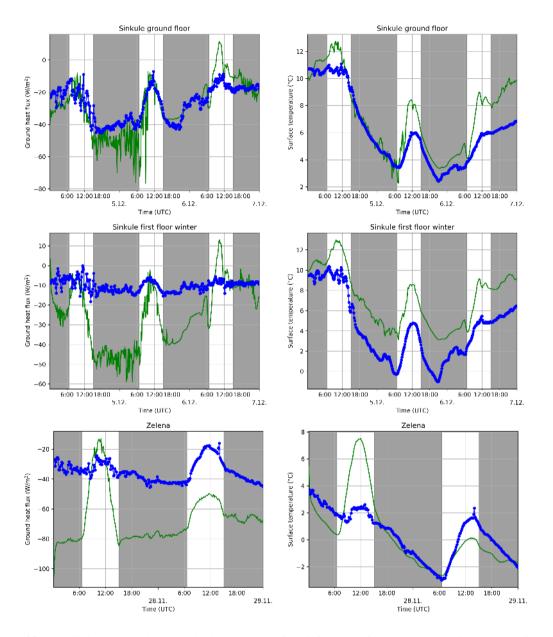
**Figure 22.** East facing wall in the street of N. A. Někrasova <u>Street</u> around location 07-1\_V (see Fig. 1 and detail location information in <u>supplements section S3.</u>). The top row shows the observed photo on 20 July 2018 at 109:37 <u>CET-UTC</u> and the 3D view of the modelled incoming SW radiation on this wall at the corresponding time step. The bottom row shows the same situation at 1110:38 <u>CETUTC</u>.



**Figure 23.** West facing wall in the street of N. A. Někrasova <u>Street</u> around location 07-2\_V (see Fig. 1 and detail location information in <u>supplements section S3.</u>). The figure shows the observed photo on 20 July 2018 at <u>1110</u>:37 <u>CET\_UTC</u> (left) and the 3D view of the modelled incoming SW radiation on this wall at the corresponding modelling time step (right).



**Figure 24.** Modelled (green) and observed (blue) wall heat flux (left) and surface temperature (right) for days 19–21 July 2018 for location Sinkule house at the ground floor wall (uppertop) and at the first floor wall (bottom). The grey areas denote the night time.



**Figure 25.** Modelled (green) and observed (blue) wall heat flux (left) and surface temperature (right) at days for 5–6 December 2018 for location Sinkule house at the ground floor (uppertop) and at the first floor (middle) and at the Zelená location Zelena at days on 27–28 November 2018 (bottom). The grey areas denote the night time.

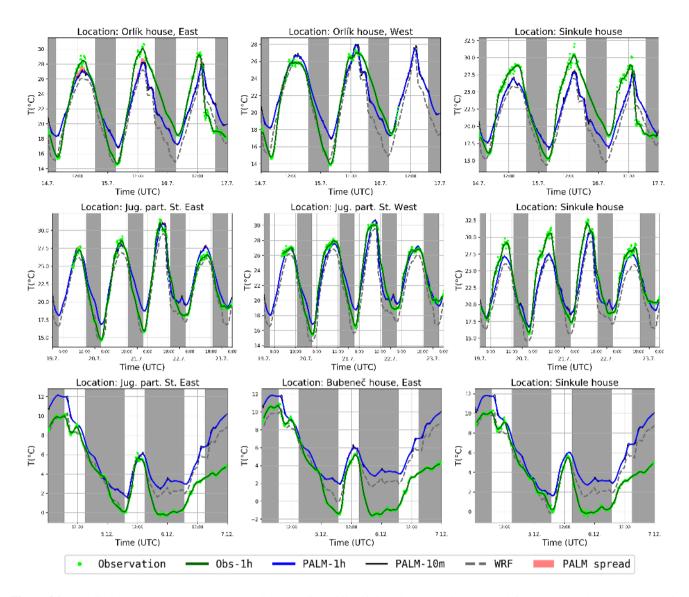


Figure 26. Modelled Street-canyon temperature at 3.9 m (solid red lineSinkule house) and observed 4.6 m (red dotsother) temperature in particular street canyon observation locations for summer e1 (top row), summer e2 (middle row) and winter e3 (bottom row) episodes. The solid red Observations are shown as 10-minute averages (green dots) and moving 1-hour averages (green linerepresents the one-hour.). PALM simulations are shown as moving 1-hour averages while the thin black (blue lineshows the original 10-minutes-), 10-minute averages . The yellow band denotes (solid black line) and the interval between the smallest and the largest 10-minute average value among the neighbouring grid points (red band). The thin dotted red grey dashed line indicates corresponding value from denotes the 1-hour averages of temperature at 2 m at the closest WRF simulationgrid point. The grey-shaded areas indicate the night time. Please note the black curve indicating the 10-minute average is mostly hidden by the blue curve indicating the one-hour average. The spatial variations in temperature are usually very small, especially in the winter case, meaning that the red curve is hidden most of the time.

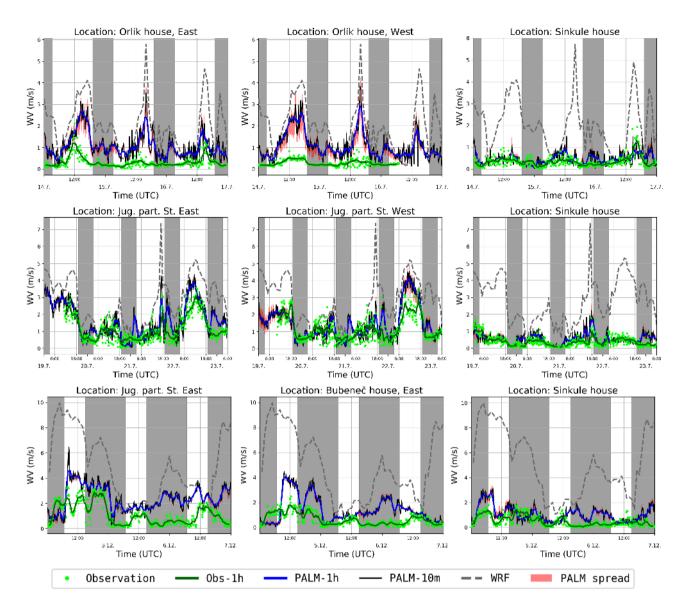
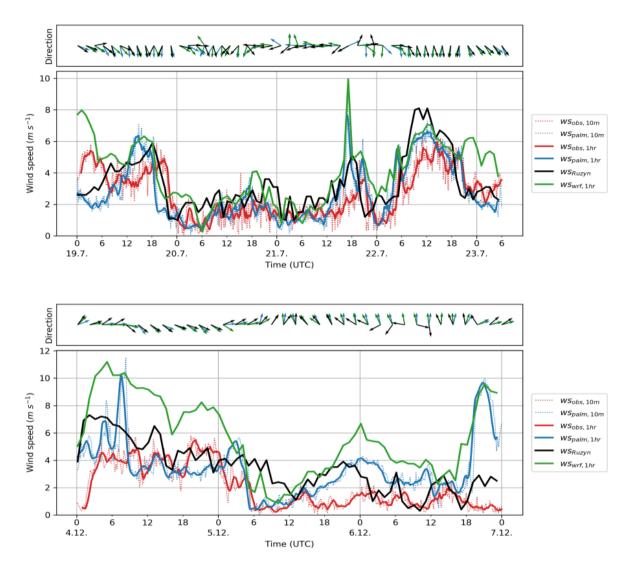
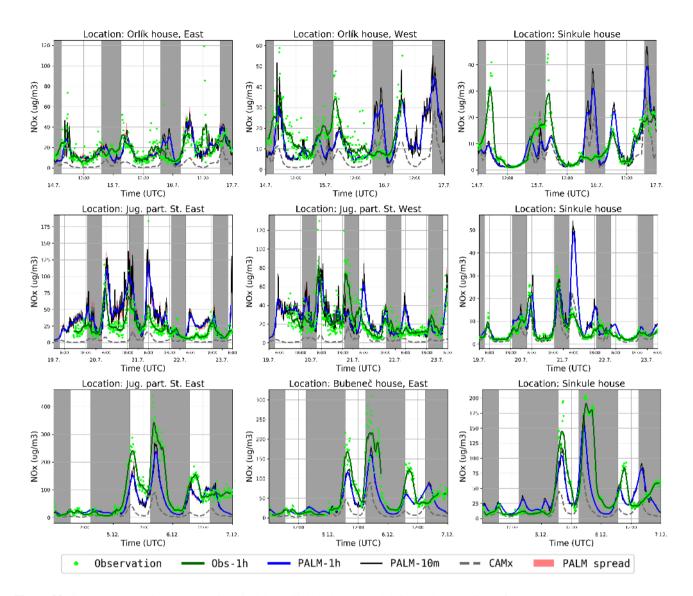


Figure 27. Modelled Street-canyon wind speed at 10 m (solid blue lineSinkule house) and observed 6.8 m (blue dotsother) wind speed in particular street canyon observation locations for summer e1 (top row), summer e2 (middle row), and winter e3 (bottom row) episodes. Solid blue Observations are shown as 10-minute averages (green dots) and moving 1-hour averages (green linerepresents one-hour). PALM simulations are shown as moving average while the darker magenta-1-hour averages (blue lineshows the original 10-minute average value averages . The light magenta band shows (solid black line) and the interval between the smallest and the largest 10-minute average value among the neighbouring grid points (red band). The thin dotted red grey dashed line indicates corresponding value from denotes 1-hour averages at 10 m at the closest WRF simulationgrid point. The grey-shaded areas indicate the night time.



**Figure 28.** Time series of wind speed and wind direction at the roof of the tallest building of the Faculty of Civil Engineering of the Czech Technical University for summer episode e2 (left) and winter episode e3 (right). The graphs show wind speed and the boxes of arrows wind direction. The red colour represents the observations, the blue colour the PALM modelled values, green colour values from the WRF model, and the black line the values from the nearest synoptic station at Praha-Ruzyně. Thin dotted lines represent 10-minute averages and the thick solid lines 1-hour moving averages of wind speed. The arrows represent 2-hour averages of wind direction. PALM model results are taken from the child domain with 2 m horizontal resolution.



**Figure 29.** Street-canyon NO<sub>X</sub> concentrations in 3.9 m (Sinkule house) and 4.6 m (other locations) for summer e1 (top row), summer e2 (middle row) and winter e3 (bottom row) episodes. Observations are shown as 10-minute averages (green dots) and moving 1-hour averages (green line). PALM simulations are shown as moving 1-hour averages (blue line), 10-minute averages (black line) and the interval between the smallest and the largest 10-minute average among the neighbouring grid points (red band). The grey dashed line denotes CAMx 1-hour concentration for the lowest level (lowest 50 m above ground) at the closest CAMx grid point. The grey-shaded areas indicate the night time.

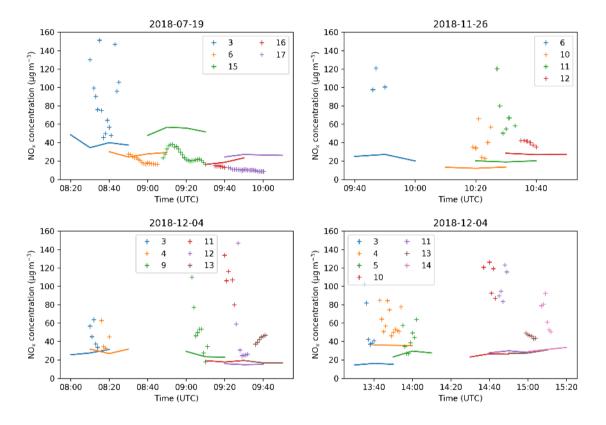


Figure 30. Mobile NO<sub>X</sub> measurements (+ marker) and modelled concentrations (solid) for 19 July 2018-morning (top left), 26 November morning (top right), 4 December 2018-morning (bottom left) and 4 December 2018-afternoon (bottom right). Numbers refer to mobile measurement locations according to Fig. 1.