

1                   Towards a model for structured mass movements: the  
2   OpenLISEM Hazard model 2.0a

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11   **Abstract**

12 Mass movements such as debris flows and landslides differ in behavior due to their material properties and  
13 internal forces. Models employ generalized multi-phase flow equations to adaptively describe these complex  
14 flow types. Such models commonly assume unstructured and fragmented flow, where internal cohesive strength  
15 is insignificant. In this work, existing work on two-phase mass movement equations are extended to include a  
16 full stress-strain relationship that allows for runout of (semi-) structured fluid-solid masses. The work provides  
17 both the three-dimensional equations and depth-averaged simplifications. The equations are implemented in a  
18 hybrid Material Point Method (MPM) which allows for efficient simulation of stress-strain relationships on  
19 discrete smooth particles. Using this framework, the developed model is compared to several flume experiments  
20 of clay blocks impacting fixed obstacles. Here, both final deposit patterns and fractures compare well to  
21 simulations. Additionally, numerical tests are performed to showcase the range of dynamical behavior produced  
22 by the model. Important processes such as fracturing, fragmentation and fluid release are captured by the model.  
23 While this provides an important step towards complete mass movement models, several new opportunities arise  
24 such as application to fragmenting mass movements and block-slides.

25

## 26 1. Introduction

27 The earths rock cycle involves sudden release and gravity-driven transport of sloping materials. These  
28 mass movements have a significant global impact in financial damage and casualties (Nadim et al., 2006;  
29 Kjekstad & Highland, 2009). Understanding the physical principles at work at their initiation and runout phase  
30 allows for better mitigation and adaptation to the hazard they induce (Corominas et al., 2014). Many varieties of  
31 gravitationally-driven mass movements have been categorized according to their material physical parameters  
32 and type of movement. Examples are slides, flows and falls consisting of soil, rocks or debris (Varnes, 1987).  
33 Major factors in determining the dynamics of mass movement runout are the composition of the moving material  
34 and the internal and external forces during initiation and runout.

35 Within the cluster of existing mass movement processes, a distinction can be made based on the  
36 cohesive of the mass during movement. Post-release, a sloping mass might be unstructured, such as mud flows,  
37 where grain-grain cohesive strength is absent. Alternatively, the mass can be fragmentative, such as strongly-  
38 deforming landslides or fragmenting of rock avalanches upon particle impacts. Finally, there are  
39 coherent/structured mass movements, such as can be the case in block-slides where internal cohesive strength  
40 can resist deformation for some period (Varnes, 1987). The general importance of the initially structured nature  
41 of mass movement material is observed for a variety of reasons. First, block slides are an important subset of  
42 mass movement types (Hayir, 2003; Beutner et al., 2008; Tang et al., 2008). This type of mass movement  
43 features some cohesive structure to the dynamic material in the movement phase. Secondly, during movement,  
44 the spatial gradients in local acceleration induce strain and stress that results in fracturing. This process, often  
45 called fragmentation in relation to structured mass movements, can be of crucial importance for mass movement  
46 dynamics (Davies & McSaveney, 2009; Delaney & Evans, 2014; Dufresne et al., 2018; Corominas et al. 2019).  
47 Lubricating effect from basal fragmentation can enhance velocities and runout distance significantly (Davies et  
48 al., 2006; Tang et al., 2009). Otherwise, fragmentation generally influences the rheology of the movement by  
49 altering grain-grain interactions (Zhou et al., 2005). The importance of structured material dynamics is further  
50 indicated by engineering studies on rock behavior and fracture models (Kaklauskas & Ghaboussi, 2001; Ngekpe  
51 et al., 2016; Dhanmeher, 2017).

52 Dynamics of geophysical flows are complex and depend on a variety of forces due to their multi-phase  
53 interactions (Hutter et al., 1996). Physically-based models attempt to describe the internal and external forces of  
54 all these mass movements in a generalized form (David & Richard, 2011; Pudasaini, 2012; Iverson & George,  
55 2014). This allows these models to be applied to a wide variety of cases, while improving predictive range. A  
56 variety of both one, two and three- dimensional sets of equations exist to describe the advection and forces that  
57 determine the dynamics of geophysical flows.

58 For unstructured (fully fragmented) mass movements, a variety of models exist relating to mohr-  
59 coulomb mixture theory. Such mass movements are described as non-Newtonian granular flows with dominant  
60 particle-particle interactions, assuming perfect mixing and continuous movement. Examples are debris flows and  
61 mudslides, while block-slides and rockslides do not fit these criteria. Within these models, the Mohr-Coulomb  
62 failure surface is described with zero cohesive strength, and only an internal friction angle (Pitman & Le, 2005).  
63 Examples that simulated a single mixed material (Rickenmann et al., 2006; O'Brien et al., 2007; Luna et al.,  
64 2012; van Asch et al., 2014). Two phase models describe both solids, fluids and their interactions and provide  
65 additional detail and generalize in important ways (Sheridan et al., 2005; Pitman & Le, 2005; Pudasaini,  
66 2012; George & Iverson, 2014; Mergili et al., 2017). Recently, a three-phase model has been developed that  
67 includes the interactions between small and larger solid phases (Pudasaini & Mergili, 2019). Typically,  
68 implemented forces include gravitational forces and, depending on the rheology of the equations, drag forces,  
69 viscous internal forces and a plasticity-criterion. The assumption of zero cohesion in the Mohr-Coulomb  
70 material is invalid for any structured mass movement. Some models do implement a non-Newtonian viscous  
71 yield stress based on depth-averaged strain estimations (Boetticher et al., 2016; Fornes et al., 2017; Pudasaini &  
72 Mergili, 2019). However, this approach lacks the process of fragmentation and internal failure.

73 For structured mass movements limited approaches are available. These movements feature some  
74 discrete inter-particle connectivity that allows the moving material to maintain a elasto-plastic structure.  
75 Examples here are block-slides rock-slides and some landslides (Aaron & Hungr, 2016). These materials can be  
76 described by a Mohr-Coulomb material with cohesive strength (Spencer, 2004). Aaron & Hungr developed a  
77 model for simulation of initially coherent rock avalanches (Aaron & Hungr, 2016) as part of DAN3D Flex.  
78 Within their approach, a rigid-block momentum analysis is used to simulate initial movement of the block. After  
79 a specified time, the block is assumed to fragment, and a granular flow model using a Voellmy-type rheology is  
80 used for further runout. Their approach thus lacks a physical basis for the fragmenting behavior. Additionally, by  
81 dissecting the runout process in two stages (discrete block and granular flow), benefits of holistic two-phase  
82 generalized runout models are lost. Finally, Greco et al. (2019) presented a runout model for cohesive granular  
83 matrix. Their approach similarly lacks a description of the fragmentation process. Thus, within current mass

84 movement models, there might be improvements available from assuming non-fragmented movement. This  
85 would allow for description of structured mass movement dynamics.

86 In this paper, a generalized mass movement model is developed to describe runout of an arbitrarily  
87 structured two-phase Mohr-Coulomb material. The model extends on recent innovations in generalized models  
88 for mohr-coulomb mixture flow (Pudasaini, 2012; Pudasaini & Mergili, 2019). The second section of this work  
89 provides the derivation of the extensive set of equations that describe structured mass movements in a  
90 generalized manner. The third section validates the developed model by comparison with results from controlled  
91 flume runout experiments. Additionally, this section shows numerical simulation examples that highlight  
92 fragmentation behavior and its influence on runout dynamics. Finally, in section four, a discussion on the  
93 potential usage of the presented model is provided together with reflection on important opportunities of  
94 improvement.

## 95 2. A set of mass movement equations incorporating internal structure

### 96 2.1 Structured mass movements

97 Gravitational mass flows are triggered when local the driving forces within a, often steep, section of a  
98 slope exceed a critical threshold. . The instability of such materials is generally understood to take place along a  
99 failure plane (Zhang et al., 2011, Stead & Wolter, 2015). Along this plane, forces exerted due to gravity and  
100 possible seismic accelerations can act as a driving force towards the downslope direction, while a normal-force  
101 on the terrain induces a resisting force (Xie et al., 2006). When internal stress exceeds a specified criteria,  
102 commonly described using Mohr-Coulomb theory, fracturing occurs, and the material becomes dynamic.  
103 Observations indicate material can initially fracture predominantly at the failure plane (Tang et al., 2009 Davies  
104 et al., 2006). Full finite-element modelling of stability confirms no fragmentation occurs at initiation, and runout  
105 can start as a structured mass (Matsui & San, 1992; Griffiths & Lane, 1999).

106 Once movement is initiated, the material is accelerated. Due to spatially non-homogeneous acceleration,  
107 either caused by a non-homogeneous terrain slope, or impact with obstacles, internal stress can build within the  
108 moving mass. The stress state can reach a point outside the yield surface, after which some form of deformation  
109 occurs (e.g. Plastic, Brittle, ductile) (Loehnert et al., 2008). In the case of rock or soil material, elastic/plastic  
110 deformation is limited and fracturing occurs at relatively low strain values (Kaklauskas & Ghaboussi, 2001;  
111 Dhanmeher., 2017). Rocks and soil additionally show predominantly brittle fracturing, where strain increments  
112 at maximum stress are small (Bieniawski, 1967; Price, 2016; Husek et al., 2016). For soil matrices, cohesive  
113 bonds between grains originate from causes such as cementing, frictional contacts and root networks (Cohen et  
114 al., 2009). Thus, the material breaks along either the grain-grain bonds or on the molecular level. In practice, this  
115 processes of fragmentation has been both observed and studied frequently. Cracking models for solids use stress-  
116 strain descriptions of continuum mechanics (Menin et al., 2009; Ngekpe et al., 2016). Fracture models frequently  
117 use Smooth Particle Hydrodynamics (SPH) since a Lagrangian, meshfree solution benefits possible fracturing  
118 behavior (Maurel & Combescure, 2008; Xu et al., 2010; Osorno & Steeb, 2017). Within the model developed  
119 below, knowledge from fracture-simulating continuum mechanical models is combined with finite element fluid  
120 dynamic models.

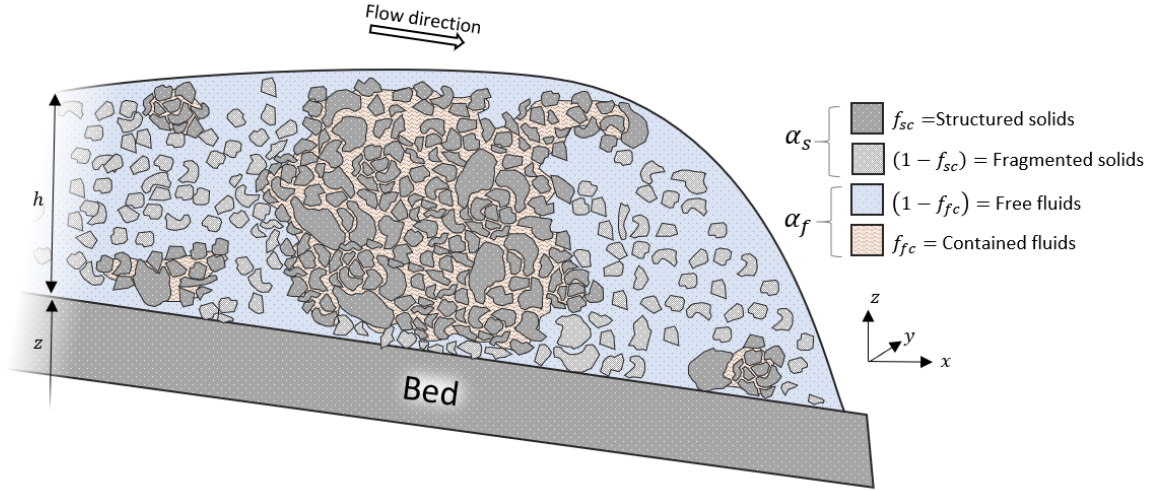
121 The mohr-coulomb mixture models on which the developed model is based, can be found in Pitman &  
122 Le (2005), Pudasaini (2012), George & Iverson, 2014 and Pudasaini & Mergili (2019). While these are  
123 commonly names debris-flow models, their validity extends beyond this typical category of mass movement.  
124 This is both apparent from model applications (Mergili et al., 2018) and theoretical considerations (Pudasaini,  
125 2012). A major cause for the usage of debris flow as a term here is the assumption of unstructured flow, which  
126 we are aiming to solve in this work.

### 127 2.2 Model description

128 We define two phases, solids and fluids, within the flow, indicated by  $s$  and  $f$  respectively. A specified  
129 fraction of solids within this mixture is at any point part of a structured matrix. This structured solid phase,  
130 indicated by  $sc$  envelops and confines a fraction of the fluids in the mixture, indicates as  $fc$ . The solids and  
131 fluids are defined in terms of the physical properties such as densities ( $\rho_f, \rho_s$ ) and volume fractions ( $\alpha_f =$   
132  $\frac{f}{f+s}, \alpha_s = \frac{s}{f+s}$ ). The confined fractions of their respective phases are indicated as  $f_{sc}$  and  $f_{fc}$  for the volume  
133 fraction of confined solids and fluids respectively (Equations 1,2 and 3).

- 134 1.  $\alpha_s + \alpha_f = 1$
- 135 2.  $\alpha_s(f_{sc} + (1 - f_{sc})) + \alpha_f(f_{fc} + (1 - f_{fc})) = 1$
- 136 3.  $(f_{sc} + (1 - f_{sc})) = (f_{fc} + (1 - f_{fc})) = 1$

137 For the solids, additionally internal friction angle ( $\phi_s$ ) and effective (volume-averaged) material size  
 138 ( $d_s$ ) are defined. We additionally define  $\alpha_c = \alpha_s + f_{fc}\alpha_f$  and  $\alpha_u = (1 - f_{fc})\alpha_f$  to indicate the solids with  
 139 confined fluids and free fluid phases respectively. These phases have a volume-averaged density  $\rho_{sc}, \rho_f$ . We let  
 140 the velocities of the unconfined fluid phase ( $\alpha_u = (1 - f_{fc})\alpha_f$ ) be defined as  $\mathbf{u}_u = (u_u, v_u)$ . We assume  
 141 velocities of the confined phases ( $\alpha_c = \alpha_s + f_{fc}\alpha_f$ ) can validly be assumed to be identical to the velocities of  
 142 the solid phase,  $\mathbf{u}_c = (u_c, v_c) = \mathbf{u}_s = (u_s, v_s)$ . A schematic depiction of the represented phases is shown in  
 143 Figure 1.



144  
 145 *Figure 1 A schematic depiction of the flow contents. Both structured and unstructured solids are*  
 146 *present. Fluids can be either free, or confined by the structured solids.*

147 A major assumption is made here concerning the velocities of both the confined and free solids (sc and  
 148 s), that have a shared averaged velocity ( $\mathbf{u}_s$ ). We deliberately limit the flow description to two phases, opposed  
 149 to the innovative work of Pudasaini & Mergili (2019) that develop a multi-mechanical three-phase model. This  
 150 choice is motivated by considerations of applicability (reducing the number of required parameters), the infancy  
 151 of three-phase flow descriptions and finally the general observations of the validity of this assumption (Ishii,  
 152 1975; Ishii & Zuber, 1979; Drew, 1983; Jakob et al, 2005; George & Iverson, 2016).

153 The movement of the flow is described initially by means of mass and momentum conservation  
 154 (Equations 4 and 5).

155 4.  $\frac{\partial \alpha_c}{\partial t} + \nabla \cdot (\alpha_c \mathbf{u}_c) = 0$

156 5.  $\frac{\partial \alpha_u}{\partial t} + \nabla \cdot (\alpha_u \mathbf{u}_u) = 0$

157 Here we add the individual forces based on the work of Pudasaini & Hutter (2003), Pitman & Le  
 158 (2005), Pudasaini (2012), Pudasaini & Fischer (2016) and Pudasaini & Mergili (2019) (Equations 6 and 7).

159 6.  $\frac{\partial}{\partial t}(\alpha_c \rho_c \mathbf{u}_c) + \nabla \cdot (\alpha_c \rho_c \mathbf{u}_c \otimes \mathbf{u}_c) = \alpha_c \rho_c \mathbf{f} - \nabla \cdot \alpha_c \mathbf{T}_c + p_c \nabla \alpha_c + \mathbf{M}_{DG} + \mathbf{M}_{vm}$

160 7.  $\frac{\partial}{\partial t}(\alpha_u \rho_f \mathbf{u}_u) + \nabla \cdot (\alpha_u \rho_f \mathbf{u}_u \otimes \mathbf{u}_u) = \alpha_u \rho_f \mathbf{f} - \nabla \cdot \alpha_u \mathbf{T}_u + p_f \nabla \alpha_u - \mathbf{M}_{DG} - \mathbf{M}_{vm}$

161 Where  $\mathbf{f}$  is the body force (among which is gravity),  $\mathbf{M}_{DG}$  is the drag force,  $\mathbf{M}_{vm}$  is the virtual mass  
 162 force and  $\mathbf{T}_c, \mathbf{T}_u$  are the stress tensors for solids with confined fluids and unconfined phases respectively. The  
 163 virtual mass force described the additional work required by differential acceleration of the phases. The drag  
 164 force describes the drag along the interfacial boundary of fluids and solids. The body force describes external  
 165 forces such as gravitational acceleration and boundary forces. Finally, the stress tensors describe the internal  
 166 forces arising from strain and viscous processes. Both the confined and unconfined phases in the mixture are  
 167 subject to stress tensors ( $\mathbf{T}_c$ , and  $\mathbf{T}_u$ ), for which the gradient acts as a momentum source. Additionally, we follow  
 168 Pudasaini (2012) and add a buoyancy force ( $p_c \nabla \alpha_c$  and  $p_f \nabla \alpha_u$ ).

### 169 **Stress Tensors, Describing internal structure**

170 Based on known two-phase mixture theory, the internal and external forces acting on the moving  
 171 material are now set up. This results in several unknowns such as the stress tensors ( $\mathbf{T}_c$  and  $\mathbf{T}_u$ , described by the

172 constitutive equation), the body force ( $\mathbf{f}$ ), the drag force ( $\mathbf{M}_{DG}$ ) and the virtual mass force ( $\mathbf{M}_{VM}$ ). This section  
 173 will first describe the derivation of the stress tensors. These describe the internal stress and viscous effects. To  
 174 describe structured movements, these require a full stress-strain relationship which is not present in earlier  
 175 generalized mass movements model. Afterwards, existing derivation of the body, drag and virtual mass force are  
 176 altered to conform the new constitutive equation.

177 Our first step in defining the momentum source terms in equations 6 and 7 is the definition of the fluid  
 178 and solid stress tensors. Current models typically follow the assumptions made by Pitman & Le (2005), who  
 179 indicate: “*Proportionality and alignment of the tangential and normal forces are imposed as a basal boundary  
 180 condition is assumed to hold throughout the layer of flowing material ... following Rankine (1857) and Terzaghi  
 181 (1936), an earth pressure relation is assumed for diagonal stress components*”. Here, the earth pressure  
 182 relationship is a vertically-averaged analytical solution for lateral forces exerted by an earth wall. Thus,  
 183 unstructured columns of moving mixtures are assumed. Here, we aim to use the full Mohr-Coulomb relations.  
 184 Describing the internal stress of soil and rock matrices is commonly achieved by elastic-plastic simulations of  
 185 the materials stress-strain relationship. Since we aim to model a full stress description, the stress tensor is equal  
 186 to the elasto-plastic stress tensor (Equation 8).

187 8.  $\mathbf{T}_c = \boldsymbol{\sigma}$

188 Where  $\boldsymbol{\sigma}$  is the elasto-plastic stress tensor for solids. The stress can be divided into the deviatoric and  
 189 non-deviatoric contributions (Equation 9). The non-deviatoric part acts normal on any plane element (in the  
 190 manner in which a hydrostatic pressure acts equal in all directions). Note that we switch to tensor notation when  
 191 describing the stress-strain relationship. Thus, superscripts ( $\alpha$  and  $\beta$ ) represent the indices of basis vectors (x, y  
 192 or z axis in Euclidian space), and obtain tensor elements. Additionally, the Einstein convention is followed  
 193 (automatic summation of non-defined repeated indices in a single term).

194 9.  $\sigma^{\alpha\beta} = s^{\alpha\beta} + \frac{1}{3}\sigma^{\gamma\gamma}\delta^{\alpha\beta}$

195 Where  $s$  is the deviatoric stress tensor and  $\delta^{\alpha\beta} = [\alpha = \beta]$  is the Kronecker delta.

196 Here, we define the elasto-plastic stress ( $\boldsymbol{\sigma}$ ) based on a generalized Hooke-type law in tensor notation  
 197 (Equation 10 and 11) where plastic strain occurs when the stress state reaches the yield criterion (Spencer, 2004;  
 198 Necas & Hiavecek, 2007; Bui et al., 2008).

199 10.  $\dot{\epsilon}_{elastic}^{\alpha\beta} = \frac{\dot{s}^{\alpha\beta}}{2G} + \frac{1-2\nu}{E}\dot{\sigma}^m\delta^{\alpha\beta}$

200 11.  $\dot{\epsilon}_{plastic}^{\alpha\beta} = \dot{\lambda}\frac{\partial g}{\partial \sigma^{\alpha\beta}}$

201 Where  $\dot{\epsilon}_{elastic}$  is the elastic strain tensor,  $\dot{\epsilon}_{plastic}$  is the plastic strain tensor,  $\dot{\sigma}^m$  is the mean stress rate  
 202 tensor,  $\nu$  is Poisson’s ratio,  $E$  is the elastic Young’s Modulus,  $G$  is the shear modulus,  $\dot{s}$  is the deviatoric shear  
 203 stress rate tensor,  $\dot{\lambda}$  is the plastic multiplier rate and  $g$  is the plastic potential function. Additionally, the strain  
 204 rate is defined from velocity gradients as equation 12.

205 12.  $\dot{\epsilon}_{total}^{\alpha\beta} = \dot{\epsilon}_{elastic}^{\alpha\beta} + \dot{\epsilon}_{plastic}^{\alpha\beta} = \frac{1}{2}\left(\frac{\partial u_c^\alpha}{\partial x^\beta} - \frac{\partial u_c^\beta}{\partial x^\alpha}\right)$

206 By solving equations 9, 10 and 11 for  $\dot{\sigma}$ , a stress-strain relationship can be obtained (Equation 13) (Bui  
 207 et al., 2008).

208 13.  $\dot{\sigma}^{\alpha\beta} = 2G\dot{\epsilon}^{\gamma\gamma}\delta^{\alpha\beta} + K\dot{\epsilon}^{\gamma\gamma}\delta^{\alpha\beta} - \dot{\lambda}\left[\left(K - \frac{2G}{3}\right)\frac{\partial g}{\partial \sigma^{mn}}\delta^{mn}\delta^{\alpha\beta} + 2G\frac{\partial g}{\partial \sigma^{\alpha\beta}}\right]$

209 Where  $\dot{\epsilon}$  is the deviatoric strain rate ( $\dot{\epsilon}^{\alpha\beta} = \dot{\epsilon}^{\gamma\gamma} - \frac{1}{3}\dot{\epsilon}^{\alpha\beta}\delta^{\alpha\beta}$ ),  $\psi$  is the dilatancy angle and  $K$  is the  
 210 elastic bulk modulus and the material parameters defined from  $E$  and  $\nu$  (Equation 14).

211 14.  $K = \frac{E}{3(1-2\nu)}, G = \frac{E}{2(1+\nu)}$

212 Fracturing or failure occurs when the stress state reaches the yield surface, after which plastic  
 213 deformation occurs. The rate of change of the plastic multiplier specifies the magnitude of plastic loading and  
 214 must ensure a new stress state conforms to the conditions of the yield criterion. By means of substituting  
 215 equation 13 in the consistency condition ( $\frac{\partial f}{\partial \sigma^{\alpha\beta}}d\sigma^{\alpha\beta} = 0$ ), the plastic multiplier rate can be defined (Equation  
 216 15) (Bui et al., 2008).

217 15.  $\dot{\lambda} = \frac{2G\dot{\epsilon}^{\alpha\beta}\frac{\partial f}{\partial \sigma^{\alpha\beta}} + \left(K - \frac{2G}{3}\right)\dot{\epsilon}^{\gamma\gamma}\frac{\partial f}{\partial \sigma^{\alpha\beta}}\sigma^{\alpha\beta}\delta^{\alpha\beta}}{2G\frac{\partial f}{\partial \sigma^{mn}}\frac{\partial g}{\partial \sigma^{mn}} + \left(K - \frac{2G}{3}\right)\frac{\partial f}{\partial \sigma^{mn}}\delta^{mn}\frac{\partial g}{\partial \sigma^{mn}}\delta^{mn}}$

218 The yield criteria specifies a surface in the stress-state space that the stress state can not pass, and at  
 219 which plastic deformation occurs. A variety of yield criteria exist, such as Mohr-Coulomb, Von Mises, Ducker-  
 220 Prager and Tresca (Spencer, 2004). Here, we employ the Ducker-Prager model fitted to Mohr-Coulomb material  
 221 parameters for its accuracy in simulating rock and soil behavior, and numerical stability (Spencer, 2004; Bui et  
 222 al., 2008) (Equation 16 and 17).

223 16.  $f(I_1, J_2) = \sqrt{J_2} + \alpha_\phi I_1 - k_c = 0$

224 17.  $g(I_1, J_2) = \sqrt{J_2} + \alpha_\phi I_1 \sin(\psi)$

225 Where  $I_1$  and  $J_2$  are tensor invariants (Equation 18 and 19).

226 18.  $I_1 = \sigma^{xx} + \sigma^{yy} + \sigma^{zz}$

227 19.  $J_2 = \frac{1}{2} s^{\alpha\beta} s^{\alpha\beta}$

228 Where the Mohr-Coulomb material parameters are used to estimate the Ducker-Prager parameters  
 229 (Equation 20).

230 20.  $\alpha_\phi = \frac{\tan(\phi)}{\sqrt{9+12 \tan^2 \phi}}, k_c = \frac{3c}{\sqrt{9+12 \tan^2 \phi}}$

231 Using the definitions of the yield surface and stress-strain relationship, combining equations 13, 15, 16  
 232 and 17, the relationship for the stress rate can be obtained (Equation 21 and 22).

233 21.  $\dot{\sigma} = 2G\dot{\epsilon}^{\alpha\beta} + K\dot{\epsilon}^{\gamma\gamma} \delta^{\alpha\beta} - \dot{\lambda} \left[ 9K \sin\psi \delta^{\alpha\beta} + \frac{G}{\sqrt{J_2}} s^{\alpha\beta} \right]$

234 22.  $\dot{\lambda} = \frac{3\alpha K \dot{\epsilon}^{\gamma\gamma} + \left(\frac{G}{\sqrt{J_2}}\right) s^{\alpha\beta} \dot{\epsilon}^{\alpha\beta}}{27\alpha_\phi K \sin\psi + G}$

235 In order to allow for the description of large deformation, the Joumann stress rate can be used, which is  
 236 a stress-rate that is independent from a frame of reference (Equation 23).

237 23.  $\hat{\sigma} = \sigma^{\alpha\gamma} \dot{\omega}^{\beta\gamma} + \sigma^{\gamma\beta} \dot{\omega}^{\alpha\gamma} + 2G\dot{\epsilon}^{\alpha\beta} + K\dot{\epsilon}^{\gamma\gamma} \delta^{\alpha\beta} - \dot{\lambda} \left[ 9K \sin\psi \delta^{\alpha\beta} + \frac{G}{\sqrt{J_2}} s^{\alpha\beta} \right]$

238 Where  $\dot{\omega}$  is the spin rate tensor, as defined by equation 24.

239 24.  $\dot{\omega}^{\alpha\beta} = \frac{1}{2} \left( \frac{\partial v^\alpha}{\partial x^\beta} - \frac{\partial v^\beta}{\partial x^\alpha} \right)$

240 Due to the strain within the confined material, the density of the confined solid phase ( $\rho_c$ ) evolves  
 241 dynamically according to equation 25.

242 25.  $\rho_c = f_{sc} \rho_s \frac{\epsilon_v^0}{\epsilon_v} + (1 - f_{sc}) \rho_s + f_{fc} \rho_f$

243 Where  $\epsilon_v$  is the total volume strain,  $\epsilon_v \approx \epsilon_1 + \epsilon_2 + \epsilon_3$ ,  $\epsilon_i$  is one of the principal components of the  
 244 strain tensor. Since we aim to simulate brittle materials, where volume strain remains relatively low, we assume  
 245 that changes in density are small compared to the original density of the material ( $\frac{\partial \rho_c}{\partial t} \ll \rho_c$ ).

## 246 Fragmentation

247 Brittle fracturing is a processes commonly understood to take place once a material internal stress has  
 248 reached the yield surface, and plastic deformation has been sufficient to pass the ultimate strength point (Maurel  
 249 & Cumescore, 2008; Husek et al., 2016). A variety of approaches to fracturing exist within the literature (Ma et  
 250 al., 2014; Osomo & Steeb, 2017). FEM models use strain-based approaches (Loehnert et al., 2008). For SPH  
 251 implementations, as will be presented in this work, distance-based approaches have provided good results  
 252 (Maurel & Cumbescure, 2008). Other works have used strain-based fracture criteria (Xu et al., 2010) .  
 253 Additionally, dynamic degradation of strength parameters have been implemented (Grady & Kipp, 1980; Vuyst  
 254 & Vignjevic, 2013; Williams, 2019). Comparisons with observed fracture behavior has indicated the predictive  
 255 value of these schemes (Xu et al., 2010; Husek et al., 2016). We combine the various approaches to best fit the  
 256 dynamical multi-phase mass movement model that is developed. Following, Grady & Kipp (1980) and we  
 257 simulate a degradation of strength parameters. Our material consists of a soil and rock matrix. We assume  
 258 fracturing occurs along the inter-granular or inter-rock contacts and bonds (see also Cohen et al., 2009). Thus,  
 259 cohesive strength is lost for any fractured contacts. We simulate degradation of cohesive strength according to a  
 260 volume strain criteria. When the stress state lies on the yield surface (the set of critical stress states within the 6-  
 261 dimensional stress-space), during plastic deformation, strain is assumed to attribute towards fracturing. A critical  
 262 volume strain is taken as material property, and the breaking of cohesive bonds occurs based on the relative

263 volume strain. Following Grady & Kipp (1980) and Vuyst & Vignjevic (2013), we assume that the degradation  
 264 behavior of the strength parameter is distributed according to a probability density distribution. Commonly, a  
 265 Weibull-distribution is used (Williams, 2019). Here, for simplicity, we use a uniform distribution of cohesive  
 266 strength between 0 and  $2c_0$ , although any other distribution can be substituted. Thus, the expression governing  
 267 cohesive strength becomes equation 26

$$268 \quad 26. \quad \frac{\partial c}{\partial t} = \begin{cases} -c_0 \frac{1}{2} \frac{\left(\frac{\epsilon_v}{\epsilon_{v0}}\right)}{\epsilon_c} & f(I_1, J_2) \geq 0, c > 0 \\ 0 & otherwise \end{cases}$$

269 Where  $c_0$  is the initial cohesive strength of the material,  $\epsilon_{v0}$  is the initial volume,  $\left(\frac{\epsilon_v}{\epsilon_{v0}}\right)$  is the fractional  
 270 volumetric strain rate,  $\epsilon_c$  is the critical fractional volume strain for fracturing.

### 271 Water partitioning

272 During the movement of the mixed mass, the solids can thus be present as a structured matrix. Within  
 273 such a matrix, a fluid volume can be contained (e.g. as originating from a ground water content in the original  
 274 landslide material). These fluids are typically described as groundwater flow following Darcy's law, which poses  
 275 a linear relationship between pressure gradients and flow velocity through a soil matrix. In our case, we assumed  
 276 the relative velocity of water flow within the granular solid matrix as very small compared to both solid  
 277 velocities and the velocities of the free fluids. As an initial condition of the material, some fraction of the water  
 278 is contained within the soil matrix ( $f_{fc}$ ). Additionally, for loss of cohesive structure within the solid phase, we  
 279 transfer the related fraction of fluids contained within that solid structure to the free fluids.

$$280 \quad 27. \quad \frac{\partial f_{fc}}{\partial t} = -\frac{\partial(1-f_{fc})}{\partial t} = \begin{cases} -f_{fc} \frac{c_0 \max(0.0, \epsilon_v)}{c \epsilon_f} & f(I_1, J_2) \geq 0, c > 0 \\ 0 & otherwise \end{cases}$$

$$281 \quad 28. \quad \frac{\partial f_{sc}}{\partial t} = -\frac{\partial(1-f_{sc})}{\partial t} = \begin{cases} -f_{sc} \frac{c_0 \max(0.0, \epsilon_v)}{c \epsilon_f} & f(I_1, J_2) \geq 0, c > 0 \\ 0 & otherwise \end{cases}$$

282 Beyond changes in  $f_{fc}$  through fracturing of structured solid materials, no dynamics are simulated for  
 283 in- or outflux of fluids from the solid-matrix. The initial volume fraction of fluids in the solid matrix defined by  
 284 ( $f_{fc}$  and  $s_{sc}$ ) remains constant throughout the simulation. The validity of this assumption can be based on the  
 285 slow typical fluid velocities in a solid matrix relative to fragmented mixed fluid-solid flow velocities (Kern,  
 286 1995; Saxton and Rawls, 2006). While the addition of evolving saturation would extend validity of the model, it  
 287 would require implementation of pretransfer-functions for evolving material properties, which is beyond the  
 288 scope of this work. An important note on the points made above is the manner in which fluids are re-partitioned  
 289 after fragmentation. All fluids in fragmented solids are released, but this does not equate to free movement of the  
 290 fluids or a disconnection from the solids that confined them. Instead, the equations continue to connect the solids  
 291 and fluids through drag, viscous and virtual mass forces. Finally, the density of the fragmented solids is assumed  
 292 to be the initially set solid density. Any strain-induced density changes are assumed small relative to the initial  
 293 solid density ( $\frac{\rho_c}{\rho_s} \ll 1$ ).

### 294 Fluid Stresses

295 The fluid stress tensor is determined by the pressure and the viscous terms (Equations 29 and 30).  
 296 Confined solids are assumed to be saturated and constant during the flow.

$$297 \quad 29. \quad \mathbf{T}_u = P_f \mathbf{I} + \boldsymbol{\tau}_f$$

$$298 \quad 30. \quad \boldsymbol{\tau}_f = \eta_f [\nabla \mathbf{u}_u + (\nabla \mathbf{u}_c)^t] - \frac{\eta_f}{\alpha_u} \mathcal{A}(\alpha_u) (\nabla \alpha_c (\mathbf{u}_u - \mathbf{u}_c) + (\mathbf{u}_c - \mathbf{u}_u) \nabla \alpha_c)$$

299 Where  $\mathbf{I}$  is the identity tensor,  $\boldsymbol{\tau}_f$  is the viscous stress tensor for fluids,  $P_f$  is the fluid pressure,  $\eta_f$  is the  
 300 dynamic viscosity of the fluids and  $\mathcal{A}$  is the mobility of the fluids at the interface with the solids that acts as a  
 301 phenomenological parameter (Pudasaini, 2012).

302 The fluid pressure acts only on the free fluids here, as the confined fluids are moved together with the  
 303 solids. In equation 30, the second term is related to the non-Newtonian viscous force induced by gradients in  
 304 solid concentration. The effect as described by Pudasaini (2012) is induced by a solid-concentration gradient. In  
 305 case of unconfined fluids and unstructured solids ( $f_{sf} = 1, f_{sc} = 1$ ). Within our flow description, we see no  
 306 direct reason to eliminate or alter this force with a variation in the fraction of confined fluids or structured solids.  
 307 We do only consider the interface between solids and free fluids as an agent that induces this effect, and  
 308 therefore the gradient of the gradient of the solids and confined fluids ( $\nabla(\alpha_s + f_{fc}\alpha_f) = \nabla\alpha_c$ ) is used instead of  
 309 the total solid phase ( $\nabla\alpha_s$ ).

## 310 Drag force and Virtual Mass

311 Our description of the drag force follows the work of Pudasaini (2012) and Pudasaini (2018), where a  
 312 generalized two-phase drag model is introduced and enhanced. We split their work into a contribution from the  
 313 fraction of structured solids ( $f_{sc}$ ) and unconfined fluids ( $1 - f_{fc}$ ) (Equation 31).

$$314 \quad 31. \quad C_{DG} = \frac{f_{sc}\alpha_c\alpha_u(\rho_c-\rho_f)g}{U_{T,c}(G(Re))+S_p}(\mathbf{u}_u - \mathbf{u}_c)|\mathbf{u}_u - \mathbf{u}_c|^{j-1} + \frac{(1-f_{sc})\alpha_c\alpha_u(\rho_s-\rho_f)g}{U_{T,uc}(\mathcal{P}\mathcal{F}(Re_p)+(1-\mathcal{P})G(Re))+S_p}(\mathbf{u}_u - \mathbf{u}_c)|\mathbf{u}_u - \mathbf{u}_c|^{j-1}$$

315 Where  $U_{T,c}$  is the terminal or settling velocity of the structures solids,  $U_{T,uc}$  is the terminal velocity of  
 316 the unconfined solids,  $\mathcal{P}$  is a factor that combines solid- and fluid like contributions to the drag force,  $G$  is the  
 317 solid-like drag contribution,  $\mathcal{F}$  is the fluid-like drag contribution and  $S_p$  is the smoothing function (Equation 32  
 318 and 34). The exponent  $j$  indicates the type of drag: linear ( $j = 0$ ) or quadratic ( $j = 1$ ).

319 Within the drag, the following functions are defined:

$$320 \quad 32. \quad F = \frac{\gamma}{180} \left(\frac{\alpha_f}{\alpha_s}\right)^3 Re_p, \quad G = \alpha_f^{M(Re_p)-1}$$

$$321 \quad 33. \quad S_p = \left(\frac{\mathcal{P}}{\alpha_c} + \frac{1-\mathcal{P}}{\alpha_u}\right)\mathcal{K}$$

$$322 \quad 34. \quad \mathcal{K} = |\alpha_c\mathbf{u}_c + \alpha_u\mathbf{u}_u| \approx 10 \text{ m s}^{-1}$$

323 Where  $M$  is a parameter that varies between 2.4 and 4.65 based on the Reynolds number (Pitman & Le,  
 324 2005). The factor  $\mathcal{P}$  that combines solid- and fluid like contributions to the drag, is dependent on the volumetric  
 325 solid content in the unconfined and unstructured materials ( $\mathcal{P} = \left(\frac{\alpha_s(1-f_{sc})}{\alpha_f(1-f_{fc})}\right)^m$  with  $m \approx 1$ ). Additionally we  
 326 assume the factor  $\mathcal{P}$ , is zero for drag originating from the structured solids. As stated by Pudasaini & Mergili  
 327 (2019) ‘‘As limiting cases:  $\mathcal{P}$  suitably models solid particles moving through a fluid’’. In our model, the drag  
 328 force acts on the unconfined fluid momentum ( $u_{uc}\alpha_f(1 - f_{fc})$ ). For interactions between unconfined fluids and  
 329 structured solids, larger blocks of solid structures are moving through fluids that contains solids of smaller size.

330 Virtual mass is similarly implemented based on the work of Pudasaini (2012) and Pudasaini & Mergili  
 331 (2019) (Equation 35). The adapted implementation considers the solids together with confined fluids to move  
 332 through a free fluid phase.

$$333 \quad 35. \quad C_{VMG} = \alpha_c\rho_u \left(\frac{1}{2}\left(\frac{1+2\alpha_c}{\alpha_u}\right)\right) \left(\left(\frac{\partial u_u}{\partial t} + u_u \cdot \nabla u_u\right) - \left(\frac{\partial u_c}{\partial t} + u_c \cdot \nabla u_c\right)\right)$$

334 Where  $C_{DG} = \frac{1}{2}\left(\frac{1+2\alpha_c}{\alpha_u}\right)$  is the drag coefficient.

## 335 Boundary conditions

336 Finally, following the work of Iverson & Denlinger (2001), Pitman & Le (2005) and Pudasaini (2012), a  
 337 boundary condition is applied to the surface elements that contact the flow (Equation 36).

$$338 \quad 36. \quad |\mathbf{S}| = N \tan(\phi)$$

339 Where  $N$  is the normal pressure on the surface element and  $\mathbf{S}$  is the shear stress.

## 340 2.3 Depth-Averaging

341 The majority of the depth-averaging in this works is analogous to the work of Pitman & Le (2005),  
 342 Pudasaini (2012) and Pudasaini & Mergili (2019). Depth-averaging through integration over the vertical extent of  
 343 the flow can be done based on several useful and often-used assumptions:  $\frac{1}{h} \int_0^h x \, dh = \bar{x}$ , for the velocities ( $u_u$   
 344 and  $u_c$ ), solid, fluid and confined fractions ( $\alpha_f$ ,  $\alpha_s$ ,  $f_{fc}$  and  $f_{sc}$ ) and material properties ( $\rho_u$ ,  $\phi$  and  $c$ ). Besides  
 345 these similarities and an identical derivation of depth-averaged continuity equations, three major differences  
 346 arise.

### 347 i) Fluid pressure

348 Previous implementations of generalized two-phase debris flow equations have commonly assumed hydrostatic  
 349 pressure ( $\frac{\partial p}{\partial z} = g^z$ ) (Pitman & Le, 2005; Pudasaini, 2012; Abe & Konagai, 2016). Here we follow this  
 350 assumption for the fluid pressure at the base and solid pressure for unstructured material (Equations 37 and 38 ).

$$351 \quad 37. \quad P_{b_{s,u}} = -(1 - \gamma)\alpha_s g^z h$$

$$352 \quad 38. \quad P_{b_u} = -g^z h$$



353 Where  $\gamma = \frac{\rho_f}{\rho_s}$  is the density ratio (not to be confused with a tensor index when used in superscript) (-).

354 However, larger blocks of structure material can have contact with the basal topography. Due to density  
 355 differences, larger blocks of solid structures are likely to move along the base (Pailhia & Pouliquen, 2009;  
 356 George & Iverson, 2014). If these blocks are saturated, water pressure propagates through the solid matrix and  
 357 hydrostatic pressure is retained. However, in cases of an unsaturated solid matrix that connects to the base,  
 358 hydrostatic pressure is not present there. We introduce a basal fluid pressure propagation factor  $\mathcal{B}(\theta_{eff}, \bar{d}_{sc}, \dots)$   
 359 which describes the fraction of fluid pressure propagated through a solid matrix (with  $\theta_{eff}$  the effective  
 360 saturation,  $\bar{d}_{sc}$  the average size of structured solid matrix blocks). This results in a basal pressure equal to  
 361 equation 39.

$$362 \quad 39. P_{bc} = -(1 - f_{sc})(1 - \gamma) \frac{(1-f_{sc})\alpha_s}{(1-f_{fc})\alpha_f} g^z h - f_{sc}(1 - \gamma)\mathcal{B} \frac{(f_{sc})\alpha_s}{(f_{fc})\alpha_f} g^z h$$

363 The basal pressure propagation factor ( $\mathcal{B}$ ) should theoretically depend, similarly to the pedotransfer  
 364 function, mostly on saturation level, as a full saturation means perfect propagation of pressure through the  
 365 mixture, and low saturation equates to minimal pressure propagation (Saxton and Rawls., 2006). Additionally it  
 366 should depend on pedotransfer functions, and the size distribution of structured solid matrices within the  
 367 mixture. For low-saturation levels, it can be assumed no fluid pressure is retained. Combined with an assumed  
 368 soil matrix height identical to the total mixture height, this results in  $\mathcal{B} = 0$ . Assuming saturation of structures  
 369 solids results in a full propagation of pressures and  $\mathcal{B} = 1$ .

### 370 ii)Stress-Strain relationship

371 Depth-averaging the stress-strain relationship in equations 22 and 23 requires a vertical solution for the  
 372 internal stress. First, we assume any non-normal vertical terms are zero (Equation 40). Commonly, Rankines  
 373 earth pressure coefficients are used to express the lateral earth pressure by assuming vertical stress to be induced  
 374 by the basal solid pressure (Equation 41 and 42) (Pitman & Le, 2005; Pudasaini, 2012; Abe & Konagai, 2016).

$$375 \quad 40. \sigma^{zx} = \sigma^{zy} = \sigma^{yz} = \sigma^{xz} = 0$$

$$376 \quad 41. \bar{\sigma}^{zz} = \frac{1}{2}P_{bs}, \sigma^{zz}|_b = P_{bs}$$

$$377 \quad 42. K_a = \frac{1-\sin(\phi)}{1+\sin(\phi)}, K_p = \frac{1+\sin(\phi)}{1-\sin(\phi)}$$

378 Here we enhance this with Bell's extension for cohesive soils (Equation 45) (Richard et al., 2017). This  
 379 lateral normal-directed stress term is added to the full stress-strain solution.

$$380 \quad 43. \bar{\sigma}_{xx} = K\sigma_{zz}|_b - 2c\sqrt{K} + \frac{1}{h} \int_0^h \sigma_{xx} dh$$

381 Finally, the gradient in pressure of the lateral interfaces between the mixture is added as a depth-  
 382 averaged acceleration term (Equation 44).

$$383 \quad 44. S_{xc} = \alpha_c \left( \frac{1}{h} \left( \frac{\partial(h\sigma^{xx})}{\partial x} + \frac{\partial(h\sigma^{yx})}{\partial y} \right) \right) + \dots$$

### 384 iii)Depth-averaging other terms

385 While the majority of terms allow for depth-averaging as proposed by Pudasaini (2012), an exception  
 386 arises. Depth-averaging of the vertical viscosity terms is required. The non-Newtonian viscous terms for the fluid  
 387 phase were derived assuming a vertical profile in the volumetric solid phase content. Here, we alter the  
 388 derivation to use this assumption only for the non-structured solids, as opposed to the structured solids where  
 389  $\frac{\partial\alpha_s}{\partial z} = 0$ .

$$390 \quad 45. \int_b^s \frac{\partial}{\partial z} \left( \frac{\partial\alpha_s}{\partial z} (u_u - u_c) \right) dz = \left[ \frac{\partial\alpha_s}{\partial z} (u_u - u_c) \right]_b^s = (\bar{u}_u - \bar{u}_c) \left[ \frac{\partial\alpha_s}{\partial z} \right]_b^s = (\bar{u}_u - \bar{u}_c) \left[ \frac{\partial\alpha_s}{\partial z} \right]_b^s =$$

$$391 \quad \frac{(\bar{u}_u - \bar{u}_c)(1-f_{sc})\zeta \bar{\alpha}_s}{h}$$

392 Where  $\zeta$  is the shape factor for the vertical distribution of solids (Pudasaini, 2012). Additionally, the  
 393 momentum balance of Pudasaini (2012) ignores any deviatoric stress ( $\tau_{xy} = 0$ ), following Savage and Hutter  
 394 (2007), and Pudasaini and Hutter (2007). Earlier this term was included by Iverson and Denlinger (2001), Pitman  
 395 and Le (2005) and Abe & Kanogai (2016). Here we include these terms since a full stress-strain relationship is  
 396 included.

### 397 Basal frictions

398 Additionally we add the Darcy-Weisbach friction, which is a Chezy-type friction law for the fluid phase  
399 that provides drag (Delestre et al., 2014). This ensures that, without solid phase, a clear fluid does lose  
400 momentum due to friction from basal shear. This was successfully done in Bout et al. (2018) and was similarly  
401 assumed in Pudasaini and Fischer (2016) for fluid basal shear stress.

$$402 \quad 46. S_f = \frac{g}{n^2} \frac{u_u |u_u|}{h^{\frac{4}{3}}}$$

403 Where  $n$  is Manning's surface roughness coefficient.

#### 404 Depth-averaged equations

405 The following set of equations is thus finally achieved for depth-averaged flow over sloping terrain (Equations  
406 47-71).

$$407 \quad 47. \frac{\partial h}{\partial t} + \frac{\partial}{\partial x} [h(\alpha_u u_u + \alpha_c u_c)] + \frac{\partial}{\partial y} [h(\alpha_u u_u + \alpha_c u_c)] = R - I$$

$$408 \quad 48. \frac{\partial \alpha_c h}{\partial t} + \frac{\partial \alpha_c h u_c}{\partial x} + \frac{\partial \alpha_c h v_c}{\partial y} = 0$$

$$409 \quad 49. \frac{\partial \alpha_u h}{\partial t} + \frac{\partial \alpha_u h u_u}{\partial x} + \frac{\partial \alpha_u h v_u}{\partial y} = R - I$$

$$410 \quad 50. \frac{\partial}{\partial t} [\alpha_c h (u_c - \gamma_c C_{VM} (u_u - u_c))] + \frac{\partial}{\partial x} [\alpha_c h (u_c^2 - \gamma_c C_{VM} (u_u^2 - u_c^2))] + \frac{\partial}{\partial y} [\alpha_c h (u_c v_c -$$

$$411 \quad \gamma_c C_{VM} (u_u v_u - u_c v_c))] = h S_{x_c}$$

$$412 \quad 51. \frac{\partial}{\partial t} [\alpha_c h (v_c - \gamma_c C_{VM} (v_u - v_c))] + \frac{\partial}{\partial x} [\alpha_c h (u_s v_s - \gamma_c C_{VM} (u_u v_u - u_c v_c))] + \frac{\partial}{\partial y} [\alpha_c h (v_c^2 -$$

$$413 \quad \gamma_c C_{VM} (v_u^2 - v_c^2))] = h S_{y_c}$$

$$414 \quad 52. \frac{\partial}{\partial t} [\alpha_u h (u_u - \frac{\alpha_c}{\alpha_u} C_{VM} (u_u - u_c))] + \frac{\partial}{\partial x} [\alpha_u h (u_u^2 - \frac{\alpha_c}{\alpha_u} C_{VM} (u_u^2 - u_c^2) + \frac{\beta_{xu} h}{2})] + \frac{\partial}{\partial y} [\alpha_u h (u_u v_u -$$

$$415 \quad \gamma_c C_{VM} (u_u v_u - u_c v_c))] = h S_{x_u} - I u_u$$

$$416 \quad 53. \frac{\partial}{\partial t} [\alpha_u h (v_u - \frac{\alpha_c}{\alpha_u} C_{VM} (v_u - v_c))] + \frac{\partial}{\partial x} [\alpha_u h (u_u v_u - \frac{\alpha_c}{\alpha_u} C_{VM} (u_u v_u - u_c v_c))] + \frac{\partial}{\partial y} [\alpha_u h (v_u^2 -$$

$$417 \quad \gamma_c C_{VM} (v_u^2 - v_c^2) + \frac{\beta_{yu} h}{2})] = h S_{y_u} - I v_u$$

$$418 \quad 54. S_{x_c} = \alpha_c \left[ g^x + \frac{1}{h} \left( \frac{\partial (h \sigma^{xx})}{\partial x} + \frac{\partial (h \sigma^{yx})}{\partial y} \right) - P_{bc} \left( \frac{u_c}{|u_c|} \tan \phi + \epsilon \frac{\partial b}{\partial x} \right) \right] - \epsilon \alpha_c \gamma_c p_{bu} \left[ \frac{\partial h}{\partial x} + \frac{\partial b}{\partial x} \right] +$$

$$419 \quad C_{DG} (u_u - u_c) |u_u - u_c|^{J-1}$$

$$420 \quad 55. S_{y_c} = \alpha_c \left[ g^y + \frac{1}{h} \left( \frac{\partial (h \sigma^{xy})}{\partial x} + \frac{\partial (h \sigma^{yy})}{\partial y} \right) - P_{bc} \left( \frac{v_s}{|\vec{u}_s|} \tan \phi + \epsilon \frac{\partial b}{\partial y} \right) \right] - \epsilon \alpha_c \gamma_c p_{bu} \left[ \frac{\partial h}{\partial y} + \frac{\partial b}{\partial y} \right] +$$

$$421 \quad C_{DG} (v_u - v_c) |v_u - v_c|^{J-1}$$

$$422 \quad 56. S_{x_u} = \alpha_u \left[ g^x - \frac{1}{2} \frac{P_{bu} h}{\alpha_u} \frac{\partial \alpha_c}{\partial x} + P_{bu} \frac{\partial b}{\partial x} - \frac{A \eta_u}{\alpha_u} \left( 2 \frac{\partial^2 u_u}{\partial x^2} + \frac{\partial^2 v_u}{\partial x y} + \frac{\partial^2 u_u}{\partial y^2} - \frac{X u_u}{\epsilon^2 h^2} \right) + \frac{A \eta_u}{\alpha_u} \left( 2 \frac{\partial}{\partial x} \left( \frac{\partial}{\partial x} (u_u - u_c) \right) + \right.$$

$$423 \quad \left. \frac{\partial}{\partial y} \left( \frac{\partial \alpha_c}{\partial x} (v_u - v_c) + \frac{\partial \alpha_u}{\partial y} (u_u - u_c) \right) \right] - \frac{A \eta_u \zeta \alpha_s (1 - f_{sc}) (u_u - u_c)}{\alpha_u h^2} - \frac{g}{n^2} \frac{u_u |u_u|}{h^{\frac{4}{3}}} \left] - \frac{1}{\gamma_c} C_{DG} (u_u -$$

$$424 \quad u_c) |\vec{u}_u - \vec{u}_c|^{J-1}$$

$$425 \quad 57. S_{y_u} = \alpha_u \left[ g^y - \frac{1}{2} \frac{P_{bu} h}{\alpha_f} \frac{\partial \alpha_c}{\partial y} + P_{bu} \frac{\partial b}{\partial y} - \frac{A \eta_u}{\alpha_u} \left( 2 \frac{\partial^2 u_f}{\partial y^2} + \frac{\partial^2 v_f}{\partial x y} + \frac{\partial^2 u_f}{\partial x^2} - \frac{X u_f}{\epsilon^2 h^2} \right) + \frac{A \eta_u}{\alpha_c} \left( 2 \frac{\partial}{\partial y} \left( \frac{\partial}{\partial y} (v_u -$$

$$426 \quad v_c) \right) + \frac{\partial}{\partial x} \left( \frac{\partial \alpha_c}{\partial y} (u_u - u_c) + \frac{\partial \alpha_c}{\partial x} (v_u - v_c) \right) \right] - \frac{A \eta_u \zeta \alpha_s (1 - f_{sc}) (v_u - v_c)}{\alpha_u h^2} - \frac{g}{n^2} \frac{v_u |u_u|}{h^{\frac{4}{3}}} \left] - \frac{1}{\gamma_c} C_{DG} (v_u -$$

$$427 \quad v_c) |\vec{u}_u - \vec{u}_c|^{J-1}$$

428

$$429 \quad 58. P_{bc} = -(1 - f_{sc})(1 - \gamma) \frac{(1 - f_{sc}) \alpha_s}{(1 - f_f) \alpha_f} g^z h - f_{sc} (1 - \gamma) \frac{(f_{sc}) \alpha_s}{(f_f) \alpha_f} g^z h$$

430

$$431 \quad 59. P_{bu} = -g^z h$$

432

$$433 \quad 60. \gamma_c = \frac{\rho_u}{\rho_c}, \gamma = \frac{\rho_f}{\rho_s}$$

$$434 \quad 61. C_{DG} = \frac{f_{sc} \alpha_c \alpha_u (\rho_c - \rho_f) g}{U_{T,c} (G(Re)) + S_p} + \frac{(1 - f_{sc}) \alpha_c \alpha_u (\rho_s - \rho_f) g}{U_{T,uc} (\mathcal{PF}(Re_p) + (1 - \mathcal{P}) G(Re)) + S_p}$$

435

436

$$\begin{aligned}
437 \quad & 62. S_p = \left( \frac{p}{\alpha_c} + \frac{1-p}{\alpha_u} \right) \mathcal{K} \\
438 \quad & 63. \mathcal{K} = |\alpha_c \mathbf{u}_c + \alpha_u \mathbf{u}_u| \\
439 \quad & 64. F = \frac{\gamma}{180} \left( \frac{\alpha_f}{\alpha_s} \right)^3 Re_p, G = \alpha_f^{M(Re_p)^{-1}}, Re_p = \frac{\rho_f d U_t}{\eta_f}, N_R = \frac{\sqrt{gLH\rho_f}}{\alpha_f \eta_f}, N_{RA} = \frac{\sqrt{gLH\rho_f}}{A\eta_f} \\
440 \quad & 65. C_{Vm} = \left( \frac{1}{2} \left( \frac{1+2\alpha_c}{\alpha_u} \right) \right) \\
441 \quad & 66. \hat{\sigma} = \sigma^{\alpha\gamma} \dot{\omega}^{\beta\gamma} + \sigma^{\gamma\beta} \dot{\omega}^{\alpha\gamma} + 2G\dot{\epsilon}^{\alpha\beta} + K\dot{\epsilon}^{\gamma\gamma} \delta^{\alpha\beta} - \dot{\lambda} \left[ 9K \sin\psi \delta^{\alpha\beta} + \frac{G}{\sqrt{J_2}} s^{\alpha\beta} \right] \\
442 \quad & 67. \dot{\lambda} = \frac{3\alpha K \dot{\epsilon}^{\gamma\gamma} + \left( \frac{G}{\sqrt{J_2}} \right) s^{\alpha\beta} \dot{\epsilon}^{\alpha\beta}}{27\alpha_\phi K \sin\psi + G} \\
443 \quad & 68. K = \frac{E}{3(1-2\nu)}, G = \frac{E}{2(1+\nu)} \\
444 \quad & 69. \sigma^{\alpha\beta} = s^{\alpha\beta} + \frac{1}{3} \sigma^{\gamma\gamma} \delta^{\alpha\beta} \\
445 \quad & 70. \dot{\epsilon}^{\alpha\beta} = \frac{1}{2} \left( \frac{\partial v^\alpha}{\partial x^\beta} - \frac{\partial v^\beta}{\partial x^\alpha} \right) \quad \dot{\omega}^{\alpha\beta} = \frac{1}{2} \left( \frac{\partial v^\alpha}{\partial x^\beta} + \frac{\partial v^\beta}{\partial x^\alpha} \right) \\
446 \quad & 71. \alpha_\phi = \frac{\tan(\phi)}{\sqrt{9+12 \tan^2 \phi}} \quad k_c = \frac{3c}{\sqrt{9+12 \tan^2 \phi}}
\end{aligned}$$

447 Where  $X$  is the shape factor for vertical shearing of the fluid ( $X \approx 3$  in Iverson & Denlinger, 2001),  $R$  is the  
448 precipitation rate and  $I$  is the infiltration rate.

#### 449 **Closing the equations**

451 Viscosity is estimated using the empirical expression from O'Brien and Julien (1985), which relates dynamic  
452 viscosity to the solid concentration of the fluid (Equation 72).

$$453 \quad 72. \eta = \alpha e^{\beta \alpha_s}$$

454 Where  $\alpha$  is the first viscosity parameter and  $\beta$  the second viscosity parameter.

455 Finally, the settling velocity of small ( $d < 100 \mu m$ ) grains is estimated by Stokes equations for a  
456 homogeneous sphere in water. For larger grains ( $> 1 mm$ ), the equation by Zanke (1977) is used (Equation 30).

$$457 \quad 73. U_T = 10 \frac{\eta}{d} \left( \sqrt{1 + \frac{0.01 \left( \frac{\rho_s - \rho_f}{\rho_f} \right) g d^3}{\frac{\eta}{\rho_f}}} - 1 \right)$$

458 In which  $U_T$  is the settling (or terminal) velocity of a solid grain,  $\eta$  is the dynamic viscosity of the fluid,  
459  $\rho_f$  is the density of the fluid,  $\rho_s$  is the density of the solids,  $d$  is the grain diameter ( $m$ )

#### 461 **2.4 Implementation in the Material Point Method numerical scheme**

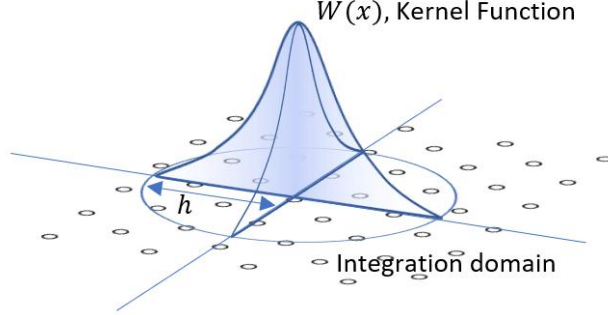
462 Implementing the presented set of equations into a numerical scheme requires considerations of that  
463 schemes limitations and strengths (Stomakhin et al., 2013). Fluid dynamics are almost exclusively solved using  
464 an Eulerian finite element solution (Delestre et al., 2014; Bout et al., 2018). The diffusive advection part of such  
465 scheme typically doesn't degrade the quality of modelling results. Solid material however is commonly  
466 simulated with higher accuracy using an Lagrangian finite element method or discrete element method (Maurel  
467 & Cumbescure, 2008; Stomakhin et al., 2013). Such schemes more easily allow for the material to maintain its  
468 physical properties during movement. Additionally, advection in these schemes does not artificially diffuse the  
469 material since the material itself is discretized, instead of the space (grid) on which the equations are solved. In  
470 our case, the material point method (MPM) provides an appropriate tool to implement the set of presented  
471 equations (Bui et al., 2008; Maurel & Cumbescure, 2008; Stomakhin et al., 2013). Numerous existing modelling  
472 studies have implemented in this method (Pastor et al., 2007; Pastor et al., 2008; Abe & Kanogai, 2016). Here,  
473 we use the MPM method to create a two-phase scheme. This allows the usage of finite elements aspects for the  
474 fluid dynamics, which are so successfully described by the that method (particularly for water in larger areas, see  
475 Bout et al., 2018).

#### 476 **Mathematical Framework**

477 The mathematic framework of smooth-particles solves differential equations using discretized volumes  
478 of mass represented by kernel functions (Libersky & Petschek, 1991; Bui et al., 2008; Stomakhin et al., 2013).  
479 Here, we use the cubic spline kernel as used by Monaghan (2000) (Equation 74).

$$480 \quad 74. \quad W(r, h) = \begin{cases} \frac{10}{7\pi h^2} \left(1 - \frac{3}{2}q^2 + \frac{3}{4}q^3\right) & 0 \leq |q| \leq 2 \\ \frac{10}{28\pi h^2} (2 - q)^3 & 1 \leq |q| < 2 \\ 0 & |q| \geq 2 \mid q < 0 \end{cases}$$

481 Where  $r$  is the distance,  $h$  is the kernel size and  $q$  is the normalized distance ( $q = \frac{r}{h}$ )



482

483 *Figure 2 Example of a kernel function used as integration domain for mathematical operations.*

484 Using this function mathematical operators can be defined. The average is calculated using a weighted  
485 sum of particle values (Equation 75) while the derivative depends on the function values and the derivative of  
486 the kernel by means of the chain rule (Equation 76) (Libersky & Petschek, 1991; Bui et al., 2008).

$$487 \quad 75. \quad \langle f(x) \rangle = \sum_{j=1}^N \frac{m_j}{\rho_j} f(x_j) W(x - x_j, h)$$

$$488 \quad 76. \quad \left\langle \frac{\partial f(x)}{\partial x} \right\rangle = \sum_{j=1}^N \frac{m_j}{\rho_j} f(x_j) \frac{\partial W_{ij}}{\partial x_i}$$

489 Where  $W_{ij} = W(x_i - x_j, h)$  is the weight of particle  $j$  to particle  $i$ ,  $r = |x_i - x_j|$  is the distance  
490 between two particles. The derivative of the weight function is defined by equation 77.

$$491 \quad 77. \quad \frac{\partial W_{ij}}{\partial x_i} = \frac{x_i - x_j}{r} \frac{\partial W_{ij}}{\partial r}$$

492 Using these tools, the momentum equations for the particles can be defined (Equations 78-84). Here, we  
493 follow Monaghan (1999) and Bui et al. (2008) for the definition of artificial numerical forces related to stability.  
494 Additionally, stress-based forces are calculated on the particle level, while other momentum source terms are  
495 solved on a Eulerian grid with spacing  $h$  (identical to the kernel size).

$$496 \quad 78. \quad \frac{dv_i^\alpha}{dt} = \frac{1}{m_i} (F_g + F_{grid}) + \sum_{j=1}^N m_j \left( \frac{\sigma_i^{\alpha\beta}}{\rho_i^2} + \frac{\sigma_j^{\alpha\beta}}{\rho_j^2} + F_{ij}^n R_{ij}^{\alpha\beta} + \Pi_{ij} \delta^{\alpha\beta} \right) \frac{\partial W_{ij}}{\partial x_i^\beta}$$

$$497 \quad 79. \quad \dot{\epsilon}^{\alpha\beta} = \frac{1}{2} \left( \sum_{j=1}^N \frac{m_j}{\rho_j} (v_j^\alpha - v_i^\alpha) \frac{\partial W_{ij}}{\partial x_i^\beta} + \sum_{j=1}^N \frac{m_j}{\rho_j} (v_j^\beta - v_i^\beta) \frac{\partial W_{ij}}{\partial x_i^\alpha} \right)$$

$$498 \quad 80. \quad \dot{\omega}^{\alpha\beta} = \frac{1}{2} \left( \sum_{j=1}^N \frac{m_j}{\rho_j} (v_j^\alpha - v_i^\alpha) \frac{\partial W_{ij}}{\partial x_i^\beta} - \sum_{j=1}^N \frac{m_j}{\rho_j} (v_j^\beta - v_i^\beta) \frac{\partial W_{ij}}{\partial x_i^\alpha} \right)$$

$$499 \quad 81. \quad \frac{d\sigma_{\alpha\beta}}{dt} = \sigma_i^{\alpha\gamma} \dot{\omega}_i^{\beta\gamma} + \sigma_i^{\gamma\beta} \dot{\omega}_i^{\alpha\gamma} + 2G_i \dot{\epsilon}_i^{\alpha\beta} + K_i \dot{\epsilon}_i^{\gamma\gamma} \delta_i^{\alpha\beta} - \lambda_i \left[ 9K_i \sin\psi_i \delta_i^{\alpha\beta} + \frac{G_i}{\sqrt{J_{2i}}} S_i^{\alpha\beta} \right]$$

$$500 \quad 82. \quad \lambda_i = \frac{3\alpha K_i \dot{\epsilon}_i^{\gamma\gamma} + \left( \frac{G_i}{\sqrt{J_{2i}}} \right) S_i^{\alpha\beta} \dot{\epsilon}_i^{\alpha\beta}}{27\alpha_\phi K_i \sin\psi_i + G_i}$$

501 Where  $i, j$  are indices indicating the particle,  $\Pi_{ij}$  is an artificial viscous force as defined by equations 83  
502 and 84 and  $F_{ij}^n R_{ij}^{\alpha\beta}$  is an artificial stress term as defined by equations 85 and 86.

$$503 \quad 83. \quad \Pi_{ij} = \begin{cases} \frac{\alpha \Pi_{sound} \phi_{ij} + \beta \Pi \phi^2}{\rho_{ij}} & v_{ij} \cdot x_{ij} < 0 \\ 0 & v_{ij} \cdot x_{ij} \geq 0 \end{cases}$$

$$504 \quad 84. \quad \phi_{ij} = \frac{h_{ij} v_{ij} x_{ij}}{|x_{ij}|^2 + 0.01 h_{ij}^2}, \quad x_{ij} = x_i - x_j, \quad v_{ij} = v_i - v_j, \quad h_{ij} = \frac{1}{2} (h_i + h_j)$$

$$505 \quad 85. \quad F_{ij}^n R_{ij}^{\alpha\beta} = \left[ \frac{W_{ij}}{W(d_0, h)} \right]^n (R_i^{\alpha\beta} + R_j^{\alpha\beta})$$

506  $86. \overline{R_i^{\gamma\gamma}} = -\frac{\epsilon_0 \overline{\sigma_i^{\gamma\gamma}}}{\rho_i^2}$

507 Where  $\epsilon_0$  is a small parameter ranging from 0 to 1,  $\alpha_{\Pi}$  and  $\beta_{\Pi}$  are constants in the artificial viscous  
 508 force (often chosen close to 1),  $u_{sound}$  is the speed of sound in the material.

509 The conversion from particles to gridded values and reversed depends on a grid basis function that  
 510 weighs the influence of particle values for a grid center. Here, a function derived from dyadic products of one-  
 511 dimensional cubic B-splines is used as was done by Steffen et al. (2008) and Stomakhin et al. (2013) (Equation  
 512 84).

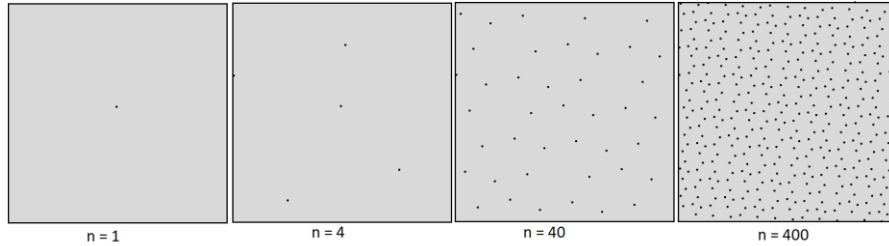
513  $87. N(\mathbf{x}) = N(x^x) * N(x^y), \quad N(x) = \begin{cases} \frac{1}{2}|x|^3 - x^2 + \frac{2}{3} & 0 \leq |x| \leq 2 \\ -\frac{1}{6}|x|^3 + x^2 - 2|x| + \frac{4}{3} & 1 \leq |x| < 2 \\ 0 & |x| \geq 2 \end{cases} \quad |x| = 0$

514 **Particle placement**

515 Particle placement is typically done in a constant pattern, as initial conditions have some constant  
 516 density. The simplest approach is a regular square or triangular network, with particles on the corners of the  
 517 network. Here, we use an approach that is more adaptable to spatially-varying initial flow height. The  $R_2$   
 518 sequence approaches, with a regular quasirandom sequence, a set of evenly distributed points within a square  
 519 (Roberts, 2020) (Equation 85).

520  $88. x_n = n\alpha \bmod 1, \quad \alpha = \left(\frac{1}{c_p}, \frac{1}{c_p^2}\right)$

521 Where  $x_n$  is the relative location of the  $n^{\text{th}}$  particle within a gridcell,  $c_p = \left(\frac{9+\sqrt{69}}{18}\right)^{\frac{1}{3}} + \left(\frac{9-\sqrt{69}}{18}\right)^{\frac{1}{3}} \approx$   
 522 1.32471795572 is the plastic constant.



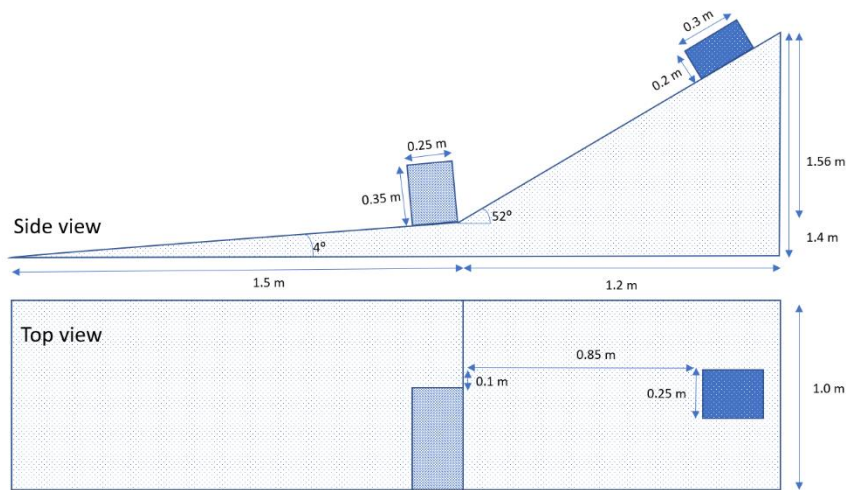
523 *Figure 3 Example particle distributions using the  $R_2$  sequence, note that, while not all particles are*  
 524 *equidistant, the method produces distributed particle patterns that adapt well to varying density.*  
 525

526 The number of particles placed for a particular flow height depends on the particle volume  $V_l$ , which is  
 527 taken as a global constant during the simulation.

528 **3. Flume Experiments**

529 **3.1 Flume Setup**

530 In order to validate the presented model, several controlled experiments were performed and reproduced  
 531 using the developed equations. The flume setup consists of a steep incline, followed by a near-flat runout plane  
 532 (Figure 3). A massive obstacle is placed on the separation point of the two planes. This blocks the path of two  
 533 fifths of the width of the moving material. For the exact dimensions of both the flume parts and the obstacle, see  
 534 figure 3.



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Figure 4 The dimensions of the flume experiment setup used in this work.

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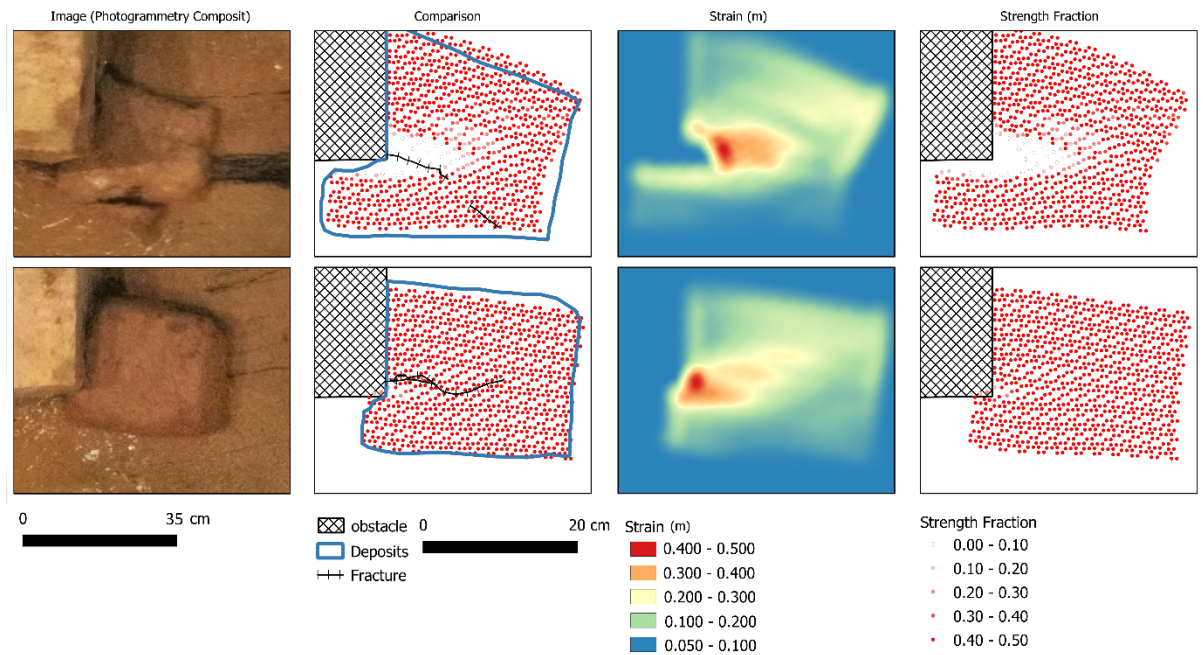
Two tests were performed whereby a cohesive granular matrix was released at the upper part of the flume setup. Both of these volumes had dimensions of 0.2x0.3x0.25 meter (height,length,width). For both of these materials, a mixture high-organic content silty-clay soils were used. The materials strength parameters were obtained using tri-axial testing (Cohesion, internal friction angle Youngs modulus and Poisson Ration. The first set of materials properties where  $c = 26.7$  kPa and  $\phi = 28^\circ$ . The second set materials properties where  $c = 18.3$  kPa and  $\phi = 27^\circ$ . For both of the events, pre-and post release elevations models were made using photogrammetry. The model was set up to replicate the situations using the measured input parameters. Numerical settings were chosen as  $\{\alpha_s = 0.5, \alpha_f = 0.5, f_{sc} = 1.0, f_{fc} = 1.0, \rho_f = 1000, \rho_s = 2400, E = 12 \cdot 10^6 Pa, K = 23 \cdot 10^6 Pa, \psi = 0, \alpha_{\Pi} = 1, \beta_{\Pi} = 1, X, \zeta, j = 2, u_{sound} = 600, dx = 10, V_l = , h = 10, n = 0.1, \alpha = 1, \beta = 10, M = 2.4, \mathcal{B} = 0, N_R = 15000, N_{RA} = 30\}$ . Calibration was performed by means of input variation. The solid fraction, and elastic and bulk modulus were varied between 20 and 200 percent of their original values with increments of 10 percent. Accuracy was assessed based on the percentage accuracy of the deposition (comparison of modelled vs observed presence of material).

550

### 3.2 Results

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Both the mapped extent of the material after flume experiments, as the simulation results are shown in figure 5. Calibrated values for the simulations are  $\{\alpha_s = 0.45, E = 21.6 \cdot 10^6 Pa, K = 13.8 \cdot 10^6 Pa\}$ .



553

554 *Figure 5 A comparison of the final deposits of the simulations and the mapped final deposits and cracks*  
 555 *within the material. From left to right: Photogrammetry mosaic, comparison of simulation results to mapped*  
 556 *flume experiment, strain, final strength fraction remaining.*

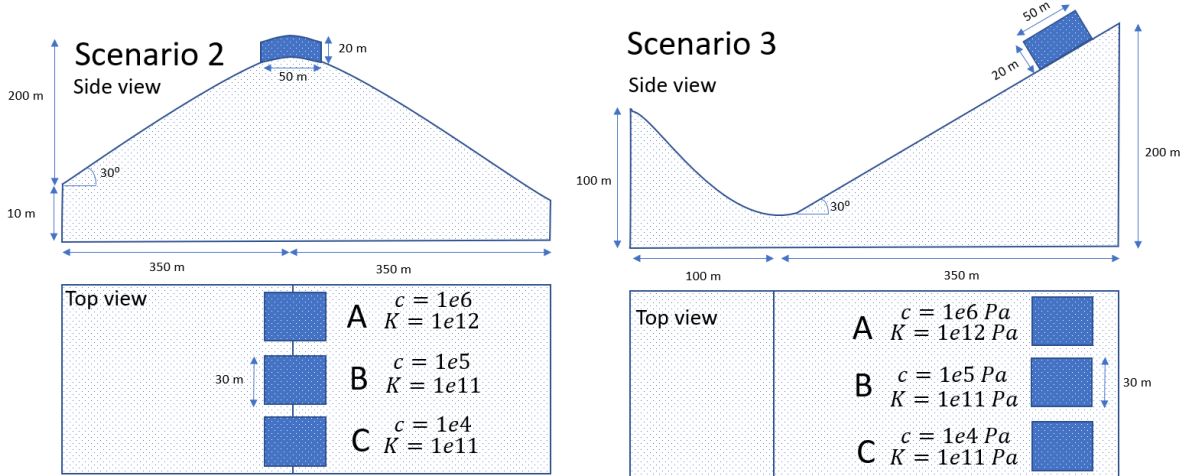
557 As soon as the block of material impacts the obstacle, stress increases as the moving objects is  
 558 deformed. This stress quickly propagates through the object. Within the scenario with lower cohesive strength,  
 559 as soon as the stress reached beyond the yield strength, degradation of strength parameters took place. In the  
 560 results, a fracture line developed along the corner of the obstacle into the length direction of the moving mass.  
 561 Eventually, this fracture developed to half the length of the moving body and severe deformation resulted. As  
 562 was observed from the tests, the first material experienced a critical fracture while the second test resulted in  
 563 moderate deformation near the impact location. Generally, the results compare well with the observed patterns,  
 564 although the exact shape of the fracture is not replicated. Several reasons might be the cause of the moderately  
 565 accurate fracture patterns. Other studies used a more controlled setup where uncertainties in applied stress and  
 566 material properties were reduced. Furthermore, the homogeneity of the material used in the tests can not  
 567 completely assumed. Realistically, minor alterations in compression used to create the clay blocks has left spatial  
 568 variation in density, cohesion and other strength parameters.

569 **4. Numerical Tests**

570 **4.1 Numerical Setup**

571 In order to further investigate some of the behaviors of the model, and highlight the novel types of mass  
 572 movement dynamics that the model implements, several numerical tests have been performed. The setup of these  
 573 tests is shown in figure 6.

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Figure 6 The dimensions of the numerical experiment setups used in this work. Setup 1 (left) and Setup 2 (right)

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Numerical settings were chosen for three different blocks with equal volume but distinct properties. Cohesive strength and the bulk modulus were varied (see figure 6). Remaining parameters were chosen as  $\{\alpha_s = 0.5, \alpha_f = 0.5, f_{sc} = 1.0, f_{fc} = 1.0, \rho_f = 1000 \text{ kgm}^{-3}, \rho_s = 2400 \text{ kgm}^{-3}, E = 1e12 \text{ Pa}, \psi = 0, \alpha_{\Pi} = 1, \beta_{\Pi} = 1, X, \zeta, j = 2, u_{sound} = 600 \text{ ms}^{-1}, dx = 10 \text{ m}, V_l, h = 10 \text{ m}, n = 0.1, \alpha = 1, \beta = 10, M = 2.4, \mathcal{B} = 0, N_R = 15000, N_{RA} = 30\}$ .

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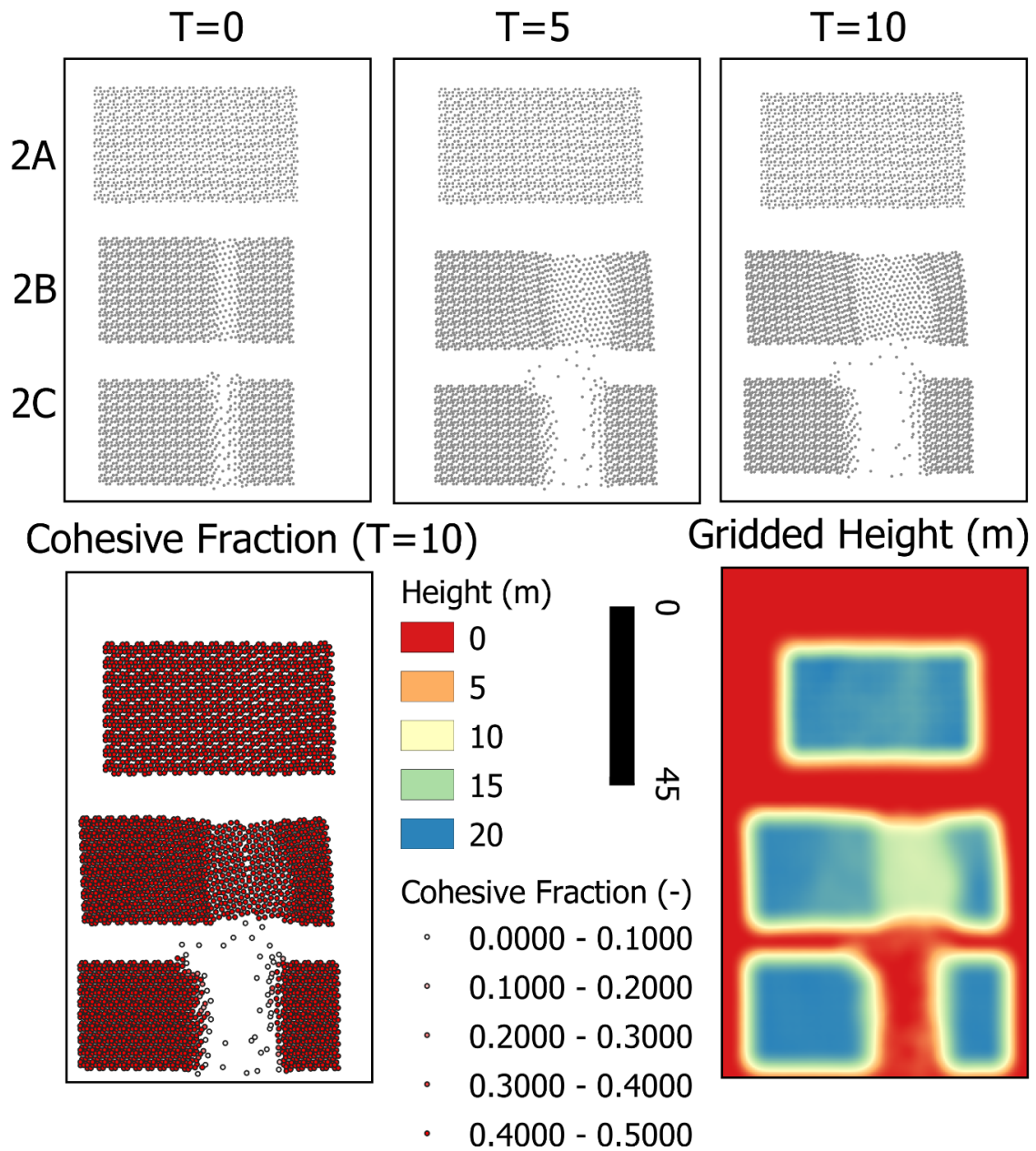
#### 4.2 Results

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Several time-slices for the described numerical scenarios are shown in figure 7 and 8.

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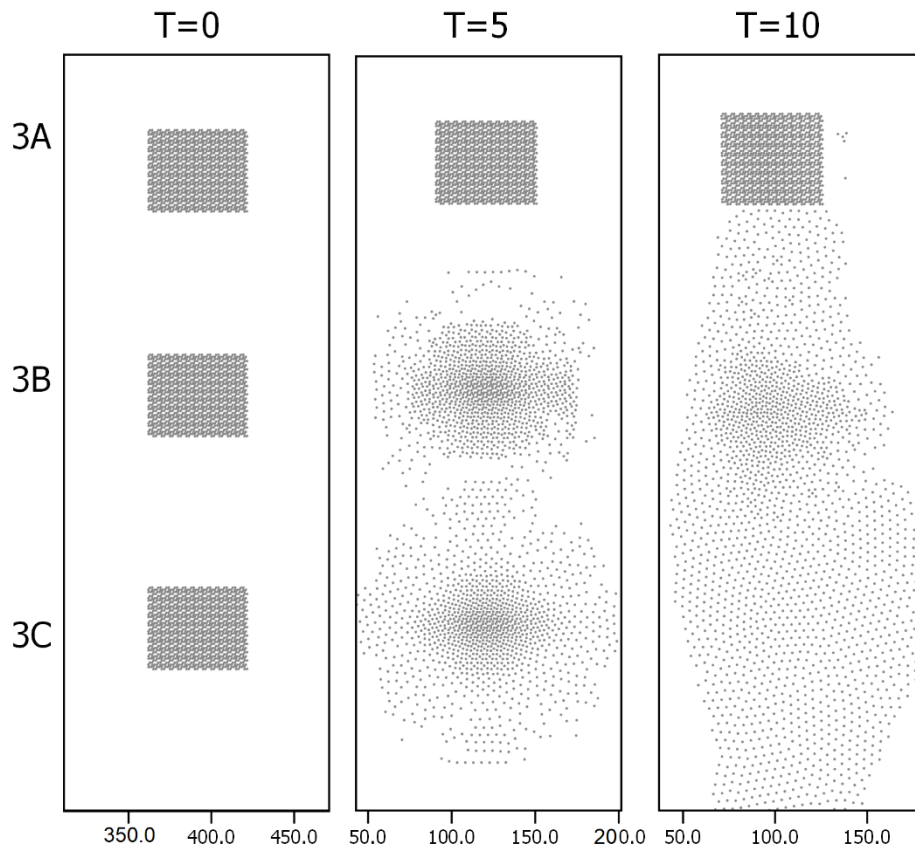


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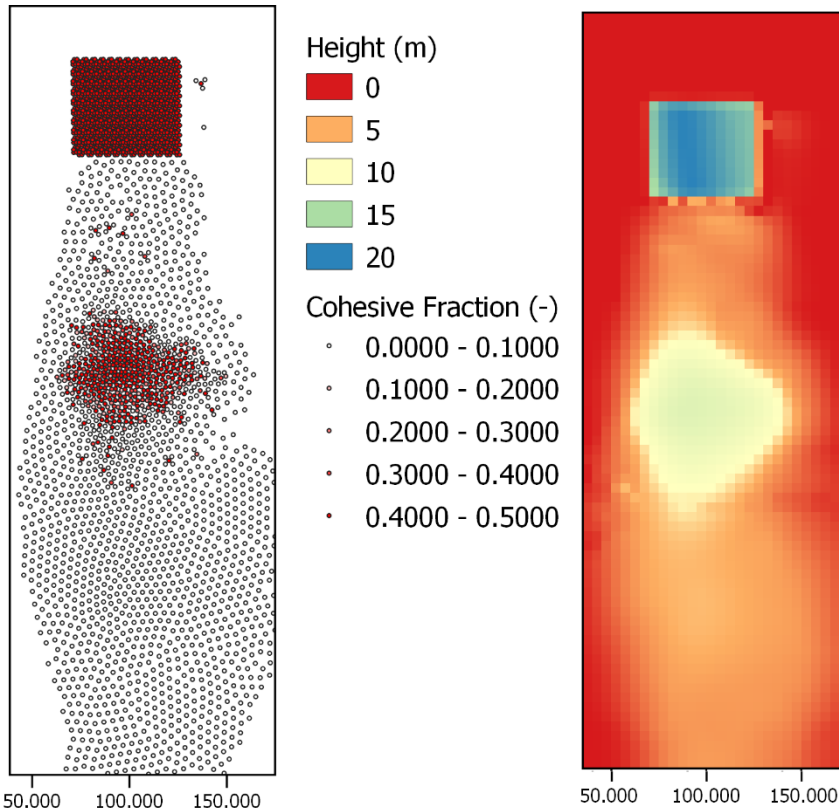
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*Figure 7 Several time-slices for numerical scenarios 2(A/B/C). See figure 6 for the dimensions and terrain setup.*



Cohesive Fraction (T=10)

Gridded Height (m)



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589

Figure 8 Several time-slices for numerical scenarios 3(A/B/C). See figure 6 for the dimensions and terrain setup.

590 Fractures develop in the mass movements based on acceleration differences and cohesive strength. For  
591 scenario 2A, the stress state does not reach beyond the yield surface, and all material is moved as a single block.  
592 Scenario 2B, which features lowered cohesive strength, fractures and the masses separate based on the  
593 acceleration caused by slopes.

594 Fracturing behavior can occur in MPM schemes due to numerical limitations inherent in the usage of a  
595 limited integration domain. Here, validation of real physically-based fracturing is present in the remaining  
596 cohesive fraction. This value only reduces in case of plastic yield, where increasing strain degrades strength  
597 parameters according to our proposed criteria. Numerical fractures would thus have a cohesive fraction of 1. In  
598 all simulated scenarios, such numerical issues were not observed.

599 Fragmentation occurs due to spatial variation in acceleration in the case of scenario 3A and 3B. For  
600 scenario 3A, the yield surface is not reached and the original structure of the mass is maintained during  
601 movement. For 3C, fragmentation is induced by lateral pressure and buoyancy forces alone. Scenario 3B  
602 experiences slight fragmentation at the edges of the mass, but predominantly fragments when reaching the  
603 valley, after which part of the material is accelerated to count to the velocity of the mass. For all the shown  
604 simulations, fragmentation does not lead to significant phase separation since virtual mass and drag forces  
605 converge the separate phase velocities to their mixture-averaged velocity. The strength of these forces partly  
606 depends on the parameters, effects of more immediate phase-separation could be studied if other parameters are  
607 used as input.

## 608 5. Discussion

609 A variety of existing landslide models simulate the behavior of lateral connected material through a  
610 non-linear, non-Newtonian viscous relationship (Boetticher et al., 2016; Fornes et al., 2017; Pudasaini &  
611 Mergili, 2019; Greco et al., 2019). These relationships include a yield stress and are usually regularized to  
612 prevent singularities from occurring. While this approach is incredibly powerful, it is fundamentally different  
613 from the work proposed here. These viscous approaches do not distinguish between elastic or plastic  
614 deformation, and typically ignore deformations if stress is insufficient. Additionally, fracturing is not  
615 implemented in these models. The approach taken in this work attempts to simulate a full stress-strain  
616 relationship with Mohr-Coulomb type yield surface. This does provide new types of behavior and can be  
617 combined with non-Newtonian viscous approaches as mentioned above. A major downside to the presented  
618 work is the steep increase in computational time required to maintain an accurate and stable simulation.  
619 Commonly, an increase of near a 100 times has been observed during the development of the presented model.

620 The presented model shows a good likeness to flume experiments and numerical tests highlight  
621 behavior that is commonly observed for landslide movements. There are however, inherent scaling issues and the  
622 material used in the flume experiments is unlikely to form larger landslide masses. The measured physical  
623 strength parameters of the material used in the flume experiments would not allow for sustained structured  
624 movement at larger scales. There is thus the need for more, real-scale, validation cases. The application of the  
625 presented type of model is most directly noticeable for block-type landslide movements that have fragmented  
626 either upon impact of some obstacle or during transition phase. Of importance here is that the moment of  
627 fragmentation is often not reported in studies on fast-moving landslides, potentially due to the complexities in  
628 knowing the details on this behavior from post-event evidence. Validation would therefore have to occur on  
629 cases where deposits are not fully fragmented, indicating that this process was ongoing during the whole  
630 movement duration. The spatial extent of initiation and deposition would then allow validation of the model.  
631 Another major opportunity for validation of the novel aspects of the model is the full three-dimensional  
632 application to landslides that were reported to have lubrication effects due to fragmentation of lower fraction of  
633 flow due to shear.

634 An important point of consideration in the development of complex multi-process generalized models is  
635 the applicability. As a detailed investigative research tool, these models provide a basic scenario of usage.  
636 However, both for research and beyond this, in applicability in disaster risk reduction decision support, the  
637 benefit drawn from these models depends on the practical requirement for parameterization and the  
638 computational demands for simulation. With an increasing complexity in the description of multi-process  
639 mechanics comes the requirement of more measured or estimated physical parameters. Inspection of the  
640 presented method shows that in principle, a minor amount of new parameters are introduced. The cohesive  
641 strength, a major focus of the model, becomes highly important depending on the type of movement being  
642 investigated. Additionally, the bulk and elastic modulus are required. These three parameters are common  
643 simulation parameters in geotechnical research and can be obtained from common tests on sampled material  
644 (Alsaman et al., 2015). Finally, the basal pressure propagation parameter ( $\mathcal{B}$ ) is introduced. However, within  
645 this work, the value of this parameter is chosen to have a constant value of one. As a result, the model does  
646 not require additional parameters, although these are relatively easy to obtain with accuracy.

647 There are a variety of aspects of the model that could be significantly improved. Here, we list several  
648 major opportunities of future research.

649 **1) Groundwater mechanics**

650 The presented model allows for the a solid or granular matrix to be present within the flow. We have  
651 assumed the flows in and out of these matrices are sufficiently small to be ignored. In reality, there is a  
652 fluid flux in and out of structured solids. This could occur both due to pressure differences as due to  
653 stress and strain of the structured solids. Implementing this kind of mechanics requires a dynamic,  
654 solid-properties dependent, soil water retention curve (Van Looy et al., 2017). An example of MPM soil  
655 mechanics with dynamic groundwater implementation can be found in Bandera et al. (2016).

656 **2) Implementing Entrainment and Deposition**

657 Current equations for entrainment (erosion with major grain-grain interactions) is limited to  
658 unstructured mixture flows (Iverson, 2012; Iverson & Ouyang, 2015; Cuomo et al., 2016; Pudasaini &  
659 Fischer, 2016). Extending these models to include a contribution from structured solids would be  
660 required to implement entrainment in the presented work.

661 **3) Separation of phases**

662 A major assumption in the presented work is that the velocities of structured solids, free solids and  
663 confined fluids are all equal. In reality, there might be separation of structured and free solids phases.  
664 Additionally, we already discussed the possibility of in-and outflux of confined fluids from the solid  
665 matrix. Recent innovations on three-phase mixture flows might be used to extend the presented work to  
666 a three, four or five-phase model by separating free solids, confined fluids or adding a Bingham-viscous  
667 solid-fluid phase (Pudasaini & Mergili, 2019). However, while this would implement an additional  
668 process, it would significantly increase complexity of the equations (in an exponential manner with  
669 relation to the number of phases) and the numerical solutions which could hinder practical applicability.

670 **4) Application to large, slow moving landslides.**

671 When confined fluids would act as a distinct phase, guided by the mechanics of water flow in granular  
672 matrix, ground water pressures and movement through the structured solids could be described. This  
673 might enable the model to do detailed deformation/groundwater simulation of large slow-moving  
674 landslides.

675 **5) Numerical Improvements**

676 Numerical techniques for particle-based discretized methods (SPH, MPM) have been proposed in the  
677 literature. A common issue is numerical fracturing of materials when particle strain increases beyond  
678 the length of the kernel function. Then, the connection between particles is lost and fracturing occurs as  
679 an artifact of the numerical method. This issue is partly solved by the artificial stress term as is also  
680 used by Bui et al. (2008). Additionally, geometric subdivide, as used by Xu et al. (2012) and Li et al.  
681 (2015), could counter these artificial fractures. Implementing this technique does require additional  
682 work to maintain mass and momentum conservation.

683 **6) Three-dimensional solutions**

684 In a variety of scenarios, the assumptions made in depth-averaged application of flow models are  
685 invalid. A common example is the impact of mass movements into lakes, or other large water bodies. In  
686 such cases, the vertical velocity and concentration variables are not well-described by their depth-  
687 averaged counterparts. Additionally, the lubrication effect of basal fragmentation of landslides due to  
688 shear can not be described without velocity-profiles and a vertical stress-solution. Full three-  
689 dimensional application would therefore have the potential to increase understanding on these important  
690 processes.

691 **5. Conclusions**

692 We have presented a novel generalized mass movement model that can describe both unstructured  
693 mixture flows and Structured movements of Mohr-Coulomb type material. The presented equations are part of  
694 the continuous development of the OpenLISEM Hazard model, an open-source tool for physically-based multi-  
695 hazard simulations. The model builds on the works of Pudasaini (2012) and Bui et al. (2008) to develop a single  
696 holistic set of equations. The model was implemented in a GPU-based Material Point Method (MPM) Code. The  
697 equations were validated on flume experiments and numerical tests, that highlight the new movement dynamics  
698 possible with the presented model. The integration of cohesive structure and a full stress-strain relationship for  
699 the structured solids allows for movement of block-type slides as a single whole. Interactions with terrain, other  
700 flow masses or obstacles lead to elastic-plastic deformation and eventually fragmentation. This type of self-  
701 alteration of flow properties is novel with mass movement models. Although the presented equations can provide  
702 additional detail for specific mass movement types, applicability of the model for real events need to be  
703 investigated as computational costs are significantly increased.

704 The presented simulation both validate the basic behavior of the model, as well as highlight the types of  
705 flow dynamics made possible by the presented equations. The models dependency of breaking to cohesive  
706 strength and internal friction angle matches the flume experiments. The numerical examples show commonly-  
707 described behavior for landslide movements. Although the simulations compare well to the flume experiments,  
708 validation is required for real-scale application to various types of mass movements. Additionally, the presented  
709 equations still lack descriptions of processes that might become important. Separating the fluid and solid phases  
710 such as done by Pudasaini & Mergili (2019), could improve flow dynamics and phase separation. With added  
711 ground-water mechanics, such as done in Bandera et al. (2016), slow-moving landslide simulations might be  
712 described.

## 713 **6. Code and Data Availability**

714 All code and data used within this work are made open-source as part of the continuous development of  
715 the OpenLISEM Hazard model under the GNU General Public Licence v3.0. The code and the data are hosted  
716 on Github (<https://github.com/bastianvandenbout/OpenLISEM-Hazard-2.0-Pre-Release>). Both binaries  
717 and a copy of the source code are also available on Sourceforge, where the manual and compilation guide can  
718 similarly be found (<https://sourceforge.net/projects/lisem/>). Finally, more information can be found at the blog  
719 (<https://blog.utwente.nl/lisem/>)

720 The software, and its user interface, are written for windows, but platform independent libraries are  
721 used and compilation might be performed on other platforms.  
722 Hardware requirements for the usage of the model are a 64-bit Operating system that can compile all required  
723 external libraries (see the manual for a full list and description). A graphical processing unit conforming to at  
724 least the OpenCL 1.2 standard and support for both OpenGL 4.2 and OpenGL/OpenCL interoperability.  
725 Additionally, an approximate 500 mb of hard drive space and 750 mb of memory must be available.

726

727 **Appendix A. List of Symbols**

728  $h$  is the flow height  
729  $s$  is the solid phase  
730  $f$  is the fluid phase  
731  $sc$  is the structured solid phase  
732  $fc$  is the confined fluid phase  
733  $\rho_f$  is the density of fluids  
734  $\rho_s$  is the density of solids  
735  $\alpha_f$  is the volumetric fluid phase fraction  
736  $\alpha_s$  is the volumetric solid phase fraction  
737  $f_{sc}$  is the fraction of solids that is structured (confining)  
738  $f_{fc}$  is the fraction of fluids that is confined  
739  $\alpha_c$  is the volumetric fraction of solids, structured solids and confined fluids  
740  $\alpha_u$  is the volumetric fraction of free fluids (unconfined phase).  
741  $\rho_{sc}$  is the volume-averaged density of the solids and confined fluids  
742  $\mathbf{u}_u$  is the velocity of the unconfined phase (free fluids)  
743  $\mathbf{u}_c$  is the velocity of the solids, confining solids and confined fluids  
744  $\mathbf{u}_s$  is the velocity of the solids  
745  $\mathbf{f}$  is the body force  
746  $\mathbf{M}_{DG}$  is the drag force  
747  $\mathbf{M}_{vm}$  is the virtual mass force  
748  $\mathbf{T}_c$  is the stress tensor for eh solids, confining solids and confined fluids  
749  $\mathbf{T}_u$  is the stress tensor for the free fluid phase  
750  $\boldsymbol{\sigma}$  is the stress tensor  
751  $\dot{\mathbf{s}}$  is the deviatoric shear stress rate tensor  
752  $\delta$  is the Kronecker delta  
753  $\dot{\epsilon}_{plastic}$  is the plastic strain rate  
754  $\dot{\epsilon}_{elastic}$  is the elastic strain rate  
755  $\lambda$  is the plastic multiplier rate  
756  $g$  is the plastic potential function  
757  $\dot{\epsilon}_{total}$  is the total strain rate  
758  $\dot{\epsilon}$  is the deviatoric strain rate  
759  $\nu$  is Poisson's ratio  
760  $E$  is the elastic Young's Modulus  
761  $G$  is the shear modulus  
762  $K$  is the Bulk elastic modulus  
763  $f(I_1, J_2)$  is the yield surface, or yield criterion  
764  $g(I_1, J_2)$  is the plastic potential function  
765  $\psi$  is the dilatancy angle  
766  $I_1$  is the first stress invariant  
767  $J_2$  is the second stress invariant  
768  $\alpha_\phi$  is the first Ducker-Prager material constant  
769  $k_c$  is the second Ducker-Prager material constant  
770  $\dot{\omega}$  is the spin rate tensor  
771  $\epsilon_{v0}$  is the initial volumetric strain  
772  $\epsilon_v$  is the volumetric strain  
773  $c_0$  is the initial cohesion  
774  $\boldsymbol{\tau}_f$  is the fluid Gauchy stress tensor  
775  $P_f$  is the fluid pressure  
776  $\eta_f$  is the fluids dynamic viscosity  
777  $\mathcal{A}$  is the mobility of the fluid at the interface  
778  $\mathcal{C}_{DG}$  is the drag coefficient  
779  $U_{T,c}$  is the settling velocity of the solids, structured solids and confined fluids  
780  $U_{T,uc}$  is the settling velocity of the unstructured solids  
781  $\mathcal{F}$  is the drag contribution from solid-like drag  
782  $\mathcal{G}$  is the drag contribution from fluid-like drag  
783  $S_p$  is the smoothing function  
784  $\mathcal{K}$  is the absolute total mass flux

785  $M(Re_p)$  is an empirical function weakly dependent on the Reynolds number  
 786  $\mathcal{P}$  the partitioning parameter for the fluid and solid like contributions to drag  
 787  $m$  is an exponent for  $\mathcal{P}$   
 788  $C_{VMG}$  is the virtual mass coefficient  
 789  $|\mathcal{S}|$  is the norm of the shear force  
 790  $N$  is the normal force on a plane element  
 791  $g$  is the gravitational acceleration  
 792  $P_{b_{s,u}}$  is the basal pressure from  
 793  $P_{b_u}$  is the basal pressure from the free fluids  
 794  $P_{b_c}$  is the basal pressure from the solids, structured solids and confined fluids  
 795  $B$  is the pressure propagation factor for structured solids  
 796  $K_a$  is the active lateral earth pressure coefficient  
 797  $K_p$  is the passive lateral earth pressure coefficient  
 798  $\zeta$  is a shape factor for the vertical gradient in solid concentration  
 799  $n$  is Mannings surface roughness coefficient  
 800  $X$  is the shape factor for the vertical fluid velocity profile  
 801  $Re_p$  is the particle Reynolds Number  
 802  $N_R$  is the Reynolds Number  
 803  $N_{RA}$  is the interfacial Reynolds Number  
 804  $H$  is the typical height of the flow  
 805  $L$  is the typical length of the flow  
 806  $\alpha$  is the first viscosity parameter  
 807  $\beta$  the second viscosity parameter  
 808  $d$  is the grain diameter  
 809  $W$  is the kernel weight function  
 810  $r$  is the distance  
 811  $h$  is the kernel width (not to be confused with the flow height)  
 812  $q$  is the normalized particle distance  
 813  $\Pi_{ij}$  is an artificial viscosity term  
 814  $F_{ij}^n R_{ij}^{\alpha\beta}$  is an artificial stress term  
 815  $\epsilon_0$  is a constant parameter for the artificial stress term  
 816  $\alpha_{\Pi}$  and  $\beta_{\Pi}$  are constants in the artificial viscous force  
 817  $u_{sound}$  is the speed of sound in the material  
 818  $N(\mathbf{x})$  is the Grid-kernel function  
 819  $c_p$  is the plastic coefficient  
 820  
 821  
 822  
 823  
 824  
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 826  
 827

828 **Appendix B. Stress Remapping**

829 If, either due to degradation of strength parameters, or building numerical errors, the state of the stress  
 830 tensor lies beyond the yield surface, a correction must be applied. We implement the correction scheme used by  
 831 Bui et al. (2008). This scheme considers two primary ways in which the stress can have an undesired state:  
 832 Tension cracking, and imperfectly plastic stress.

833 **Tension Cracking**

834 In the case of tension cracking, the stress state has moved beyond the apex of the yield surface, as  
 835 described by Chen & Mizuno (1990). The employed solution in this case is to re-map the stress tensor along the  
 836  $I_1$  axis to be at this apex. The apex is provided by the yield function (Equation 89)

837 89.  $-\alpha_\phi I_1 + k_c < 0$

838 To solve for this condition, the non-deviatoric stress state is increased (since  $I_1 - \frac{k_c}{\alpha_\phi}$  is negative) to lie  
 839 perpendicular to the apex point on the  $I_1$  axis (Equation ).

840 90.  $\widetilde{\sigma}^{\gamma\gamma} = r s^{\gamma\gamma} - \frac{1}{3} \left( I_1 - \frac{k_c}{\alpha_\phi} \right)$

841 **Imperfect Plastic Stress**

842 Imperfect plastic stress described the state where the stress tensor lies above the apex, but beyond the  
 843 yield criterion, thus have more stress than supported by the failure criteria that is set. This criteria is simply the  
 844 yield surface itself (Equation 91).

845 91.  $-\alpha_\phi I_1 + k_c < \sqrt{J_2}$

846 For this state, re-mapping is done by scaling of the  $J_2$  value (Equations 92, 93 and 94).

847 92.  $r = \frac{-\alpha_\phi I_1 + k_c}{\sqrt{J_2}}$

848 93.  $\widetilde{\sigma}^{\gamma\gamma} = r s^{\gamma\gamma} + \frac{1}{3} I_1$

849 94.  $\widetilde{\sigma}^{\bar{x}\bar{y}} = r s^{xy}, \widetilde{\sigma}^{\bar{x}\bar{y}} = r s^{xz}, \widetilde{\sigma}^{\bar{x}\bar{y}} = r s^{yz}$

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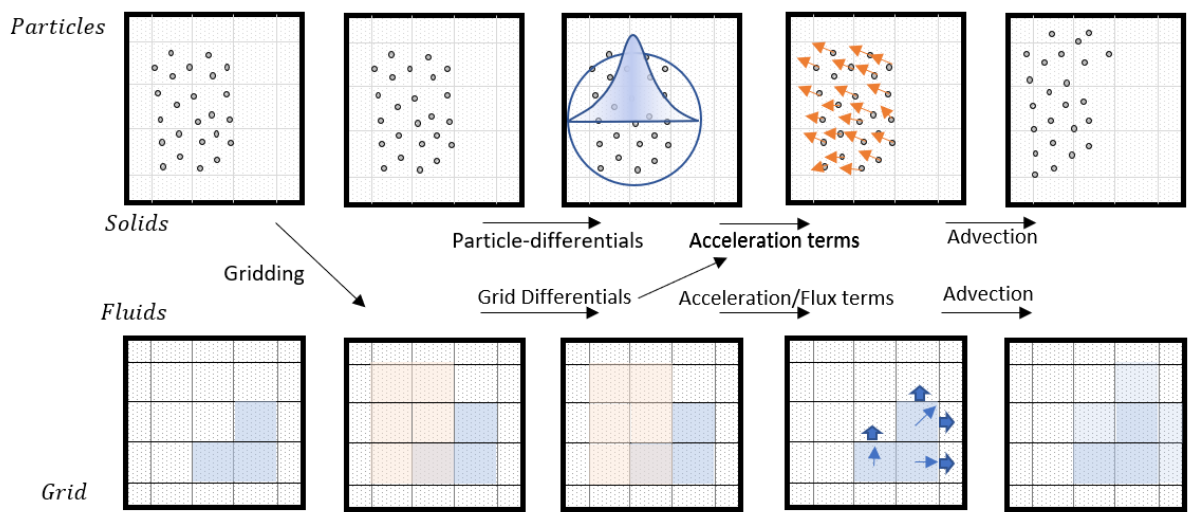


853 **Appendix C. Software Implementation**

854 The model presented in this article is part of the continued development of the OpenLISEM modelling  
 855 tools. The most recent set of equations of implemented in the open-source alpha version of OpenLISEM Hazard  
 856 2. Here, we describe the details of the implementation of the model into software.

857 **Hybrid MPM**

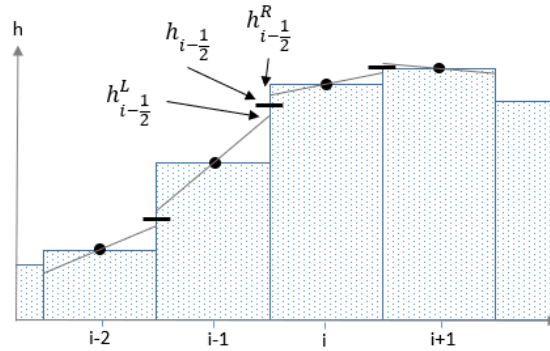
858 We utilize the MPM framework to be able to discretize part of the equations on a Eulerian regular grid,  
 859 and part of the equations on the Lagrangian particles. Our distinct take on this method is the representation of the  
 860 fluid phase completely as a finite element solution, while solids are simulated as discrete particle volumes. This  
 861 allows the model to use the major benefits that are present when depth-averaged fluid flow is simulated in a grid.  
 862 Both numerical efficiency, and high-accuracy coupling with hydrology are lacking in particle methods. For the  
 863 solid phase, non-dissipative advection, fracturing and stiffness is a major benefit of the MPM approach. Since  
 864 our model assumed confined fluids share their velocity with the solids, we advect the confined fluids as part of  
 865 the particles. Total fluid volume is then calculated from the free fluids in the finite element data, and the gridded  
 866 particle data. A flowchart of the software setup is provided in figure 6.



867  
 868 *Figure 9 The sub-steps taken by the software to complete a single step of numerical integration.*

869 **Finite element solution**

870 We use a regular cartesian grid to describe the modelling domain. Terrain and cell-boundary based  
 871 variables are re-produced using the MUSCL piecewise linear reconstruction (Delestre et al., 2014). For each cell-  
 872 boundary, a left and right estimation of acceleration terms, velocity updates and new discharges is made. The left  
 873 estimates use left-reconstructed variables while the other uses right-reconstructed variables. The final average  
 874 flux through the boundary determines actual mass and momentum transfer. Local acceleration is averaged from  
 875 the right estimate of the left boundary and left estimate of the right boundary. An additional benefit of the used  
 876 scheme is the automatic estimation of continuous and discontinuous terrain. The piecewise linear reconstructions  
 877 do not guarantee smooth terrain, for sharp locally variable terrain, pressure terms from vertical walls arise that  
 878 block momentum. These terms allow for better estimation of momentum loss by barriers, but can be turned off if  
 879 required for the simulated scenario.

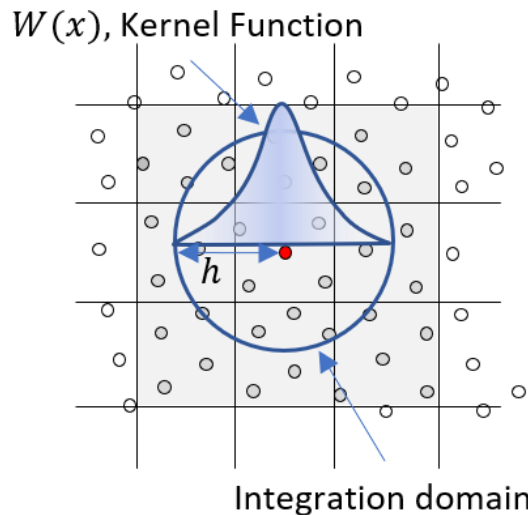


880

881 *Figure 10 Piecewise linear reconstruction is used by the MUSCL scheme to estimate values of flow*  
 882 *heights, velocities and terrain at cell-boundaries.*

883 **GPU acceleration using OpenCL/OpenGL**

884 In order to create a more efficient setup, both the finite element and particle interactions are performed  
 885 on the GPU. We utilize the OpenCL API to compile kernels written in c-style language. These kernels are  
 886 compiled at the start of the simulation, and thereby allow for easy customization by users. While the usage of  
 887 OpenCL 1.1 forces the usage of single precision floating point numbers, it allows for a wider range of GPU types  
 888 to be supported. Finite element solutions on the GPU are straightforward, as maps are a basic data storage type  
 889 for graphical processing units. Particles are stored as single-precision floating point arrays. Within the  
 890 framework of MPM, iteration of particles within a kernel is required for each timestep and particle. This  
 891 effectively means  $O(n^2)$  operations are required. Significant efficiency improvements are obtained by pre-  
 892 calculation sorting. Particles are sorted based on their location within the finite element grid. Based on the id of  
 893 the gridcell, a bitonic mergesort is performed. This sorting algorithm works seamlessly on parallel architecture  
 894 and operates as  $O(n \log^2(n))$  (Batcher, 1968). The then, a raster is allocated to store the first indexed occurrence  
 895 within the sorted list of particles of that gridcell. Since the kernel used for the presented work extends at most to  
 896 a full width of two gridcells, we must iterate over all particles present in 9 neighboring grid cells.



897

898 *Figure 11 By limiting the kernel with and sorting particles before calculation, only the distance of*  
 899 *particles in neighboring cells need to be checked, significantly reducing computational load, particularly for*  
 900 *larger datasets.*

901 A final benefit to the usage of OpenCL is direct access to simulation variables for visualization in  
 902 OpenGL using the OpenGL/OpenCL interoperability functionality. The built-in viewing window of OpenLISEM  
 903 Hazard 2.0 alpha directly uses the data to draw both particles, shapefiles and grid data using customizable  
 904 shaders written in the OpenGL shader language.



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