Improving the LPJmL4-SPITFIRE vegetation-fire model for South America using satellite data

Markus Drüke^{1,2}, Matthias Forkel³, Werner von Bloh¹, Boris Sakschewski¹, Manoel Cardoso⁴, Mercedes Bustamante⁵, Jürgen Kurths^{1,2}, and Kirsten Thonicke¹

Correspondence: Markus Drüke (drueke@pik-potsdam.de)

Abstract. Vegetation fires influence global vegetation distribution, ecosystem functioning, and global carbon cycling. Specifically in South America, changes in fire occurrence together with land use change accelerate ecosystem fragmentation and increase the vulnerability of tropical forests and savannas to climate change. Dynamic Global Vegetation Models (DGVMs) are valuable tools to estimate the effects of fire on ecosystem functioning and carbon cycling under future climate changes. However, most fire-enabled DGVMs have problems in capturing the magnitude, spatial patterns, and temporal dynamics of burnt area as observed by satellites. As fire is controlled by the interplay of weather conditions, vegetation properties and human activities, fire modules in DGVMs can be improved in various aspects. In this study we focus on improving the controls of climate and hence fuel moisture content on fire danger in the LPJmL4-SPITFIRE DGVM in South America and especially for the Brazilian fire-prone biomes Caatinga and Cerrado. We therefore test two alternative model formulations (standard Nesterov index and a newly implemented water vapor pressure deficit) for climate effects on fire danger within a formal model-data integration setup where we estimate model parameters against satellite data sets of burnt area (GFED4) and above ground biomass of trees. Our results show that the optimized model improves the representation of spatial patterns and the seasonal to interannual dynamics of burnt area especially in the Cerrado/Caatinga region. In addition, the model improves the simulation of above-ground biomass and the spatial distribution of plant functional types (PFTs). We obtained the best results by using the water vapor pressure deficit (VPD) for the calculation of fire danger. The VPD includes, in comparison to the Nesterov index, a representation of the air humidity and the vegetation density. This work shows the successful application of a systematic model-data integration setup, as well as the integration of a new fire danger formulation, in order to optimize a process-based fire-enabled DGVM. It further highlights the potential of this approach to achieve a new level of accuracy in comprehensive global fire modelling and prediction.

¹Potsdam Institute for Climate Impact Research, Telegraphenberg A 31, Potsdam, Germany

²Humboldt Universität zu Berlin, Unter den Linden 6, 10099 Berlin, Deutschland

³ TU Wien, Department of Geodesy and Geoinformation, Gusshausstr. 27-29, 1040 Vienna, Austria

⁴Instituto Nacional de Pesquisas Espaciais, Av. dos Astronautas, 1.758 - Jardim da Granja, São José dos Campos - SP, 12227-010, Brazil

⁵Instituto de Ciências Biologicas, Universidade de Brasília, Campus Universitário Darcy Ribeiro - Asa Norte, 70910-900 Brasília, Brazil

1 Introduction

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Fire in the Earth system is an important disturbance leading to many changes in the vegetation and has substantial impact on biodiversity, human health and ecosystems (Langmann et al., 2009). Fire is responsible for ca. 2 Pg of carbon emissions, which constitutes 20 % of global carbon emissions (Giglio et al., 2013; Werf et al., 2010). Fire-induced aerosol emissions and land surface changes modify evapotranspiration and surface albedo and have therefore a crucial impact on global climate (van der Werf et al., 2008; Yue and Unger, 2018). Despite a tendency for globally declining burnt area (Andela et al., 2017; Forkel et al., 2019b), more frequent and intense drought-periods lead to increasing fire-prone weather and surface conditions worldwide and therefore fire danger (Jolly et al., 2015). Growing fire danger along with land-use change are increasing ecosystem's vulnerability, which could in turn shift entire regions into a less vegetated state (Silvério et al., 2013). To account for these effects, it is extremely important to include well performing fire modules in Dynamic Global Vegetation Models (DGVMs).

Especially in South America, tropical forests, woodlands and other ecosystems are vulnerable to increasing fire danger and land use change (Cochrane and Laurance, 2008). This study focuses on the fire behavior in central-northern South America and especially on the Brazilian biomes Caatinga and Cerrado, which are the most fire-prone regions in South America (Fig. 1). Together with the Amazon rainforest they form an area of very high biodiversity and have a large impact on the global carboncycle and the regional water cycle (Lahsen et al., 2016). Compared to the Amazon, the Cerrado and Caatinga are both less

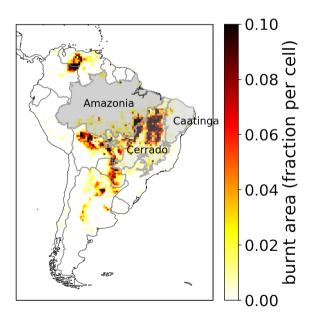


Figure 1. Overview of the mean annual burnt area in Brazil from 2005-2015 (Werf et al., 2017; Giglio et al., 2013) and the biomes Amazonia, Cerrado and Caatinga (IBGE, 2019; Harvard, 2019)

densely vegetated and drier biomes, but with very different vegetation and precipitation dynamics. The Cerrado is a savannalike biome with a mixture of shrubs, high grasses and dry forest parts. With a precipitation of ca. 1500 mm per year the Cerrado does experience a rainy season. The Caatinga, on the other hand, has a semi-arid climate with irregular rainfall between 500 and 750 mm per year, mostly within only a few months of the year. The vegetation is heterogeneous and characterized by deciduous dry forest and shrubs (Alvares et al., 2013; Prado, 2003). The different vegetation types of the Caatinga and the Cerrado lead to different fire spread, fire intensity, fire resistance and fire mortality properties. While within the Cerrado fire is a frequent event and the plants are mostly adapted to it (70 % of burnt area in Brazil is within the Cerrado, Moreira de Araújo et al., 2012), the Caatinga has a lower fire intensity and fire spread due to a lower biomass, available for fuel. Such variability in the vegetation and dead fuel composition, within and between biomes, poses a challenge to global fire models to correctly simulate observed fire patterns for a variety of biomes. Both, the Caatinga and the Cerrado depend on a strict equilibrium of fire-vegetation-climate feedbacks (Lasslop et al., 2016), which is threatened to be disturbed by human impact through climate change and land use change (Beuchle et al., 2015). While the Amazon is the focus of various national and international conservation efforts and at least by law well protected, the Cerrado is currently over-exploited by the agribusiness and its importance for regional climate, biodiversity and the water cycle is often neglected (Lahsen et al., 2016). In particular the disturbance of increasing fire regimes by climate change and land-use change might accelerate biome degradation. These effects on the Cerrado might also impact the Amazon rainforest by shifting the position of the savanna-forest biome boundary towards forest, putting the functioning of the Amazon rain forest at risk (Chambers and Artaxo, 2017). Parts of the Cerrado are also itself vulnerable to increasing fire regimes, and might shift to a less vegetated state, similar to the Caatinga (Hoffmann et al., 2000). To model these feedback processes and to study the range of biome-stability under certain drought-induced perturbations, a realistic fire representation in climate and vegetation models is essential. However, modelling fire behavior of the Brazilian Cerrado and Caatinga presents a huge challenge.

The fire occurrence depends on many interconnected parameters as humidity, precipitation, temperature, ignition sources (lightning and human) and windspeed, but also on fuel load, fuel moisture and the adaption of plant traits to fire (Keeley et al., 2011), which makes the development of fire models a complex task (Forkel et al., 2019a; Hantson et al., 2016; Lasslop et al., 2015; Krawchuk and Moritz, 2011; Jolly et al., 2015). Global fire modelling is done either by empirical models (e.g. Thonicke et al., 2001; Knorr et al., 2016; Forkel et al., 2017) or by process-based models (e.g. Venevsky et al., 2002; Thonicke et al., 2010). Empirical fire models are simplified statistical representations of fire processes and are based on empirical relationships between variables (e.g. soil moisture and fire occurrence). Process-based fire models attempt to simulate fire via explicit process-based relations: Fire ignitions are calculated by taking into account lightning flashes as natural sources and human ignitions. The chance of an ignition to become a spreading fire is then determined by the fire danger index. Sophisticated fire models calculate the rate of spread by taking into account wind speed and then translate these results into an area burnt, fuel consumption and fire carbon emissions (e.g. Thonicke et al. (2010); see Hantson et al. (2016) for an overview of global fire models).

Weather conditions control the moisture content of fuels and the danger of fire to ignite and spread. Hence the simulation of fire danger plays an important role to simulate the occurrence of fire within global process-based fire models (Pechony and Shindell, 2009). Temperature, precipitation, humidity and vegetation-related variables are often used to compute fire weather indices and hence to estimate the risk of ignitions to become a spreading fire (Chuvieco et al., 2010). Various fire weather indices are used within operational fire danger assessment systems (e.g. Canadian Fire Weather index, FWI (Wagner et al.,

1987), the Keetch Byram Drought Index (Keetch and Byram, 1968), the Angström Fire Danger Index (Arpaci et al., 2013), and the Nesterov index (Venevsky et al., 2002)). However, regional studies show that fire weather indices tend to have different predictive performances for fire occurrence (Arpaci et al., 2013). Hence, the performance of different fire weather indices should be ideally tested in order to accurately represent fire danger in DGVMs. However, not all fire weather indices can be easily adapted for global fire models because they require input variables that are not available within a DGVM framework. Hence a fire danger index for a DGVM should be as complex as necessary but still relatively easy to implement. As a result, the relatively simple Nesterov index has been widely used within global fire models (Venevsky et al., 2002; Thonicke et al., 2010).

Here, we aim to improve the simulated occurrence of fire (i.e. burnt area) in the LPJmL4-SPITFIRE model for South America and in particular for the fire-prone biomes Cerrado and Caatinga. We aim to evaluate the performance of two alternative fire danger indices within SPITFIRE based on the already implemented Nesterov index (Venevsky et al., 2002) and the newly implemented water vapor pressure deficit (VPD thereafter, Pechony and Shindell, 2009; Ray et al., 2005). Furthermore, we apply a formal model-data integration framework (LPJmLmdi, Forkel et al., 2014) to estimate model parameters that control fire danger, fire behavior, fire resistance and mortality against satellite-based data sets of burnt area and above-ground biomass (Fig. 2). Our approach is likely to improve the representation of spatial-temporal variations in fire behavior in different biomes to enable a much better modelling of the impact of climate change on fire-vegetation interactions in the current century.

2 Materials and Methods

2.1 The coupled vegetation-fire model LPJmL4-SPITFIRE

2.1.1 LPJmL 4.0

The LPJmL 4.0 model (Lund-Potsdam-Jena managed Land, Schaphoff et al., 2018a, b), is a well established and validated process-based DGVM, which globally simulates the surface energy balance, water fluxes, carbon fluxes and stocks, and natural and managed vegetation from climate and soil input data. LPJmL simulates global vegetation distribution as the fractional coverage of plant functional types (PFT), which is called foliage projective cover (FPC), and managed land as fractional coverage of crop functional types (CFT). The establishment and survival of different PFTs is regulated through bioclimatic limits and effects of heat, productivity and fire on plant mortality. Therefore, it enables LPJmL to investigate feedbacks, for example between vegetation and fire. In standard settings, which are also used here, the model operates on the grid of $0.5^{\circ} \times 0.5^{\circ}$ latitude-longitude with a spinup time of 5000 years, repeating the first 30 years of the given climate data set. Since its original implementation by Sitch et al. (2003), LPJmL has been improved by a representation of the water balance (Gerten et al., 2004), a representation of the agriculture (Bondeau et al., 2007), and new modules for fire (Thonicke et al., 2010), permafrost (Schaphoff et al., 2013) and phenology (Forkel et al., 2014).

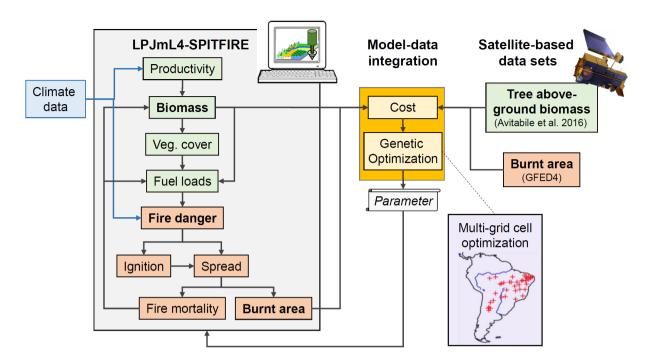


Figure 2. Schematic overview of the model-data integration approach to estimate parameters of LPJmL4-SPITFIRE against satellite-based data sets of burnt area and above-ground biomass

2.1.2 SPITFIRE

SPITFIRE (SPread and InTensity of FIRE, Thonicke et al., 2010) is a process-based fire module, used in various vegetation models (e.g. Lasslop et al., 2014; Yue and Unger, 2018), including LPJmL4. We describe here its main features, which are published in Thonicke et al. (2010). SPITFIRE calculates fire disturbance by simulating the ignition, the danger, the spread and the effects of fire separately. As ignition sources SPITFIRE considers human ignition and lightning flashes. Human ignitions $(n_{h,ig})$ are calculated as a function of population density:

$$n_{h,iq} = P_D \cdot k(P_D) \cdot a(N_D)/100, \tag{1}$$

where

$$k(P_D) = 30.0 \cdot \exp(p_h \cdot \sqrt{P_D}). \tag{2}$$

 P_D is the human population density (individuals km⁻²) and $a(N_D)$ (ignitions individual⁻¹ day⁻¹) describes the inclination of humans to cause fire ignitions (Eq. 3 and 4 in Thonicke et al., 2010). p_h is a parameter, which is set to -0.5 in Thonicke et al. (2010). This relationship assumes that human ignitions are lowest on very low populated regions and on high populated regions through a higher level of urbanization and landscape fragmentation. The ignition is highest for a medium-small population density. Lightning-caused ignitions are prescribed by lightning data from the OTD/LIS Gridded Climatological data set

(Christian et al., 2003), assuming that 20 % of the flashes reach the ground and 4 % of cloud-to-ground strikes can start a fire. In the study area of South America human ignitions are by far the most dominant ignition source, due to missing lightning in the dry season.

Fire danger is by default computed by using the Nesterov index which accounts for the maximum and dew point temperatures as well as scaling factors for different PFTs on a daily time step. In the following section, we describe the calculation of the fire danger indices in detail (Sect. 2.2). Fire duration t_{fire} (min) is calculated as a function of the fire danger index, assuming that fires burn longer under a high fire danger:

$$t_{fire} = \frac{241}{1 + 240 \cdot \exp(p_t \cdot FDI)},\tag{3}$$

where p_t is set to -11.06 in Thonicke et al. (2010). The maximum fire duration per day is 240 minutes.

The calculation of the forward rate of spread $ROS_{f,surface}$ (m min⁻¹) is based on the Rothermel equations (Rothermel, 1972; Pyne et al., 1996; Wilson, 1982):

$$ROS_{f,surface} = \frac{I_R \cdot \xi \cdot (1 + \Phi_w)}{\rho_b \cdot \epsilon \cdot Q_{ig}},\tag{4}$$

where I_R is the reaction intensity, ξ the propagation flux ratio, Φ_w a multiplier that accounts for the effect of wind, ϵ the effective heating number, Q_{ig} the heat of pre-ignition and ρ_b the fuel bulk density (Eq. 9 in Thonicke et al., 2010). ρ_b (kg m⁻³) is a PFT-dependent parameter and describes the density of the fuel, which is available for burning. It is weighted over the different fuel classes. Hence, a changing PFT distribution has an impact on $ROS_{f,surface}$.

The simulated fire ignitions, fire danger and fire spread are then used to calculate the burnt area, fire carbon emissions, and plant mortality. Plant mortality depends on the scorch height (SH) and the probability of mortality due to crown damage $P_m(CK)$. SH describes the height of the flame at which canopy scorching occurs. It increases with the 2/3 power of the surface intensity

20 $I_{surface}$:

$$SH = F \cdot I_{surface}^{0.667},\tag{5}$$

where F is a PFT-dependent parameter. Assuming a cylindrical crown, the proportion CK affected by fire is calculated as:

$$CK = \frac{SH - H + CL}{CL},\tag{6}$$

where H is the height of the average woody PFT and CL the crown length. The probability of mortality $P_m(CK)$ due to crown damage is then calculated by:

$$P_m(CK) = rCK \cdot CK^p, \tag{7}$$

where rCK is a PFT depended resistance factor between 0 and 1, and p in the range of 3 to 4. Disturbance by fire mortality has a large impact on the vegetation dynamics, which are calculated within LPJmL. SPITFIRE further includes a surface intensity threshold (10^6 , fraction burnt area per grid cell), which describes the threshold of the possible area burnt below which the surface intensity is set to zero and hence burnt area, emissions and fuel consumption is set to zero.

SPITFIRE considers anthropogenic effects on fire by taking into account human ignitions but does not account for fire suppression. Only wildfires occurring in natural vegetation are simulated. Fire on managed land like agriculture or pasture areas is not implemented, which has to be taken into account if simulated burnt area is compared with satellite observation.

Furthermore, we introduced a small technical change in the LPJmL4 interaction with SPITFIRE compared to the original SPITFIRE implementation: In the version 4.0 of LPJmL the fire litter routine calculates the leaf and litter carbon pools in a daily time step. Since the LPJmL tree allocation works at a yearly time step, this implementation leads to an incorrect LPJmL4-SPITFIRE interaction. We now split the fire-litter routine into two parts; the first one allocates burnt matter into the litter at a daily time step without recalculating the pools and the second one calculates the leaf and root carbon pools at a yearly time step.

10 2.2 Fire danger indices

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The fire danger index (FDI) is a key parameter within process-based fire models such as SPITFIRE. The FDI determines the probability and the intensity of a spreading fire, which impacts fire behavior.

2.2.1 Nesterov index-based fire danger index (FDI_{NI})

The fire danger index within SPITFIRE is based on the daily (d) calculated Nesterov Index NI(d) (Venevsky et al., 2002), which is widely used in numeric fire simulations. The NI is a cumulative function of daily maximum temperature $T_{max}(d)(^{\circ}C)$ and dew-point temperature $T_{dew}(d)(^{\circ}C)$ and set to zero at a precipitation ≥ 3 mm or a temperature $\leq 4 ^{\circ}C$:

$$NI(d) = \sum T_{max}(d) \cdot (T_{max}(d) - T_{dew}(d)), \tag{8}$$

$$T_{dew} = T_{min}(d) - 4. (9)$$

The resulting fire danger index has been calculated as in Schaphoff et al. (2018a) (slightly different compared to Thonicke et al. (2010)) by taking into account the NI as measure for weather conditions and a PFT-dependent scaling factor α_{NI_i} :

$$FDI_{NI} = max \left(0, 1 - \frac{1}{m_e} exp \left(-\frac{\sum \alpha_{NI_i}}{n} \cdot NI \right) \right), \tag{10}$$

where n is the number of PFTs and m_e the moisture of extinction, which is a PFT-dependent parameter and is weighted over the litter amount. We will use the scaling factors α_{NI_i} in the parameter optimization (Sect. 2.4).

2.2.2 Vapor pressure deficit-based fire danger index (FDI_{VPD})

We implemented a new fire danger index, based on the water vapor pressure deficit (VPD). The VPD describes the difference of the saturation water vapor pressure e_s and the actual water vapor pressure in the air. For the parameterization of the VPD

we used an approach based on Pechony and Shindell (2009):

$$VPD \propto 10^{Z(T)} (1 - RH/100),$$
 (11)

where T is the air temperature, RH the relative humidity and Z the Goff-Gratch equation (Goff and Gratch, 1946) to calculate the saturation vapor pressure. The flammability F at time step t for each grid cell can then be expressed as:

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$$F(t) = 10^{Z(T(t))} \left(1 - \frac{RH(t)}{100}\right) VD(t)e^{-c_R R(t)},$$
 (12)

where VD is the vegetation density, R the total precipitation in mm/day and c_R is a constant factor ($c_R = 2$ day/mm). Here we used the simulated FPC from LPJmL4 as a proxy for the VD. The soil is a natural buffer for drought periods and heavy rainfall events. In the Nesterov index this was taken into account by the cumulative nature of this index. Since the VPD-based fire danger index is not cumulative, this buffering effect is taken into account by taking the monthly mean of the precipitation. In doing so we avoid unrealistic high flammability fluctuations in time steps with isolated events of very low or very high precipitation (R).

Based on this implementation in SPITFIRE, the resulting FDI was much smaller than the original FDI_{NI} . Hence, we scaled the VPD up with a PFT-dependent scaling factor α_{VPD_i} , weighted over the corresponding FPC:

$$FDI_{VPD} = \frac{\sum \alpha_{VPD_i} \cdot FPC_i}{\sum FPC_i} \cdot F(t). \tag{13}$$

 α_{VPD_i} for the FDI_{VPD} was not included in (Pechony and Shindell, 2009), but is important in order to allow different fire responses for different tree and grass types. We will use the scaling factors α_{VPD_i} in the parameter optimization (Sect. 2.4). In comparison to the NI, the FDI_{VPD} requires more climate variables as input as it uses relative humidity and vegetation cover as additional fire-relevant variables. Vegetation cover has a direct link to fire risk by providing the number of available fuel for burning. According to many studies (e.g. Ray et al., 2005; Sedano and Randerson, 2014; Seager et al., 2015) the FDI_{VPD} is a very accurate fire danger index with a high correlation with fire occurrence, while still being relatively easy to implement in a global fire model.

The general behavior of the two indices as modelled by LPJmL in dependence of relative humidity and temperature is shown in Fig. 3. The Nesterov index shows a strong but very localized maximum for high temperatures and a small humidity. Hence a spreading fire is only possible in a very small climate range (here ca. from 25° Celsius and a relative humidity smaller than 0.5). The VPD on the other hand shows a less pronounced maximum but a medium fire danger also for wetter and colder regions.

The slope of towards lower VPD values is also smaller compared to the Nesterov index. Especially in regions with temperatures colder than 20°C and a relative humidity smaller than ca. 0.6 a fire is still possible. This might increase the area in which fires can occur compared to the Nesterov index, which could be an important improvement, enabling SPITFIRE to simulate more fire in wetter and colder regions. The calculated VPD and NI values shown in Fig. 3 are based on a LPJmL-SPITFIRE run, and

thus the influence of vegetation distribution on both fire danger indices.

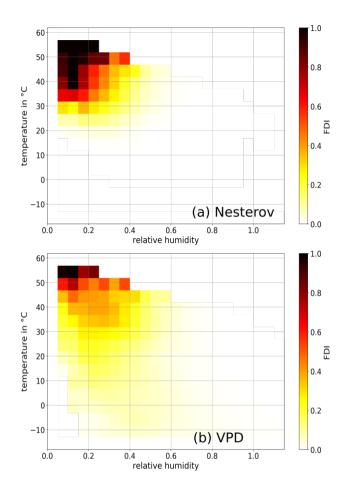


Figure 3. Dependence of the simulated fire danger index on monthly mean relative humidity and temperature for (a) the Nesterov-based index and (b) the VPD-based index. Both indices were calculated with monthly data for the years 2000-2010.

2.3 Model input data

LPJmL4-SPITFIRE requires input data on daily air temperature, precipitation, long-wave and shortwave downward radiation, wind and specific humidity, which are taken from the NOAH Global Land Assimilation System (GLDAS, Rodell et al., 2004). The data has a spatial resolution of $0.25^{\circ} \times 0.25^{\circ}$ and the time step is 3 hours. We regridded and aggregated the data set to the LPJmL resolution of $0.5^{\circ} \times 0.5^{\circ}$ and to a daily time step. We used the GLDAS 2.0 for the years 1948-1999 and the version GLDAS 2.1 for the years 2000-2017. GLDAS 2.1 uses multiple satellite- and ground-based observational data as well as advanced land surface modelling and data assimilation techniques. GLDAS 2.0 is forced entirely with the Princeton meteorological forcing data (Civil and Environmental Engineering/Princeton University, 2006). Because LPJmL4 requires at least 30 years of climate data for its spin-up (Sect. 2.1.1), the time span covered by GLDAS 2.1 is too short. To run the model, we used both climate data sets, but used the years 2003-2013 from GLDAS 2.1 for the optimization and 2005-2015 for the evaluation

period.

Furthermore, LPJmL4-SPITFIRE is forced with gridded constant soil texture (Nachtergaele et al., 2009) and annual information on land use from Fader et al. (2010). Atmospheric CO_2 concentrations are used from Mauna Loa station (Quéré et al., 2015) and applied globally. The population density is taken from Goldewijk et al. (2011) and the lightning flashes are taken from the OTD/LIS satellite product (Christian et al., 2003).

2.4 Model optimization

To estimate parameters of LPJmL4-SPITFIRE, we aimed to calibrate model results against satellite observations of burnt area (GFED4: Giglio et al., 2013; Werf et al., 2017). However, as fire occurrence and spread impact and depend on vegetation productivity, hence fuel load, we wanted to ensure to not over-fit LPJmL4 against burnt area but to additionally achieve a realistic vegetation distribution. Therefore, we additionally included a satellite-derived data set on above-ground biomass of trees (AGB, Avitabile et al., 2016) in the optimization. We combined burnt area and AGB with the corresponding model outputs within a joint cost function and applied a genetic optimization algorithm to estimate model parameters (Fig. 2). The implementation of the genetic optimization algorithm (GENetic Optimization Using Derivatives (GENOUD), Mebane and Sekhon, 2011) for LPJmL is described in Forkel et al. (2014). The used cost function is based on the Kling-Gupta efficiency (KGE), which is the Euclidean distance in a three-dimensional space of model performance measures that account for the bias, ratio of variance and correlation between simulations and the observations. Gupta et al. (2009) showed that the KGE performs in an optimization setup is better than, e.g., the Nash-Sutcliffe efficiency (and hence MSE). We extended the KGE by defining it for multiple data sets d (i.e. burnt area and AGB):

$$Cost = \sqrt{\sum_{d=1}^{N} \left(\frac{s'_d}{o'_d} - 1\right)^2 + \left(\frac{\sigma_{s,d}}{\sigma_{o,d}} - 1\right)^2 + (r(s_d, o_d) - 1)^2}$$
(14)

where \acute{s} and \acute{o} are mean values (bias component) over space (i.e. different grid-cells) and time (e.g. months) of simulations s and the observations o, respectively. σ_s and σ_o are variances (variance component) and r is the Pearson correlation coefficient over space and time. The optimization was performed for 40 grid-cells in South America to represent a variety of fire regimes (Fig. 2). We selected the grid cells manually to cover active fire regions (either in the model or in the evaluation data), specifically in the Cerrado and Caatinga. We selected a high density of grid cells in the Caatinga region to improve the very poor model performance in this region. To make sure that the model performance in the Caatinga and Cerrado was not achieved at the cost of a poor performance in other areas, we also additionally selected some cells in areas where initial fire modeling gave good results, as well as in areas where minimal or no fire occurs (central Brazilian Amazon). After inspection of the results, minor adjustments were made and the selection of the grid cells was modified to account for neglected regions (which showed worsening of the model performance). These initial analysis actually demonstrate that the choice of grid cells is important for the model optimization and requires the development of a more thorough selection method in future model optimization applications.

Several parameters of LPJmL4-SPITFIRE were included in the optimization that cover different fire processes (see Tab. 2): ignition (human ignition parameter p_h , Eq. 2), fire danger (scaling factors FDI (α_{NI_i} and α_{VPD_i}), Eq. 10 and 13), fire spread (fire duration p_t , Eq. 3), fuel bulk density (ρ_b , Eq. 4), surface intensity threshold and fire effects (scorch height parameter F,

Eq. 5; crown mortality parameter rCK, Eq. 7). While p_t , p_h and the surface intensity threshold are global parameters (for all PFTs), the others were optimized for each PFT separately. Since we focus here on tropical South America, we used tropical broadleaved evergreen (TrBE), tropical broadleaved raingreen (TrBR) and tropical herbaceous (TrH) PFTs for the optimization.

In genetic optimization algorithms, each model parameter is called an individual with a corresponding fitness, which represents the cost of the model against the observations. At the beginning of the optimization process, the GENOUD algorithm creates a generation of individuals based on random sampling of parameter sets within the prescribed parameter ranges. After the calculation of the cost of all individuals of the first generation, a next generation is generated by cloning the best individuals, by mutating the genes or by crossing different individuals (Mebane and Sekhon, 2011). This results, after some generations, in a set of individuals with highest fitness, i.e. parameter sets with minimized cost. To find an optimum parameter set also used the BFGS gradient search algorithm (named after the authors Broyden, 1970; Fletcher, 1970; Goldfarb, 1970; Shanno, 1970) within the GENOUD algorithm. An optimized parameter set of the BFGS algorithm is used as individual in the next generation. We were applying the GENOUD algorithm with 20 generations and a population size of 800 individuals per generation, which corresponds to 16000 single model runs. We decided on this amount of iterations, because the cost kept almost constant in the last iterations and the parameter values did not change to the 6th digit, beyond which changes are not really relevant for model applications. During the optimization we ran the model parallel for each grid cell (40 grid cells and CPU's, 3.2GHz) and had a total optimization time of ca. 24 hours.

The comparison of the two presented fire danger indices is the main objective of this study. Hence the optimization of the PFT-dependent FDI scaling factors α_{NI_i} and α_{VPD_i} is crucial and obligatory for the VPD because of no prior values. Accordingly, we conducted two different optimization experiments using LPJmLmdi: First, using a FDI based on the VPD (VPD_{optim} hearafter) and secondly using the a FDI based on the NI (NI_{optim} hereafter). Both resulting parameter sets were then used for LPJmL4 runs and were compared to the unoptimized original model version using the NI (NI_{orig} hereafter) and various evaluation data sets.

25 2.5 Evaluation data

We used burnt area from the global fire emission database (GFED4; Giglio et al., 2013; Werf et al., 2017), in the model optimization and to evaluate model results. The global data set is available at a resolution of $0.25^{\circ} \times 0.25^{\circ}$ in a monthly time step from 1997 until 2016. The GFED burnt area product is based on the 500 m Collection 5.1 MODIS direct broadcast (DB) burnt area product (MCD64A1, after 2001). We used data for the years 2003-2013 in the optimization in order to not include potential inconsistencies between the GLDAS 2.0 and 2.1 climate data sets or between burnt area observations within GFED that originate from different satellite sensors. The GFED product comes with a stratification of burnt area by land cover from the MODIS land cover map in the resolution of 500 m (Giglio et al., 2013). As LPJmL does not simulate fire on managed lands, we excluded burnt area on cropland classes from model-data comparisons. Due to lack of data we however did not account for the proportion of pastures. To constrain the simulated vegetation distribution, we used the AGB data set from Avitabile et al.

(2016). This data set is approximately representative for the late 2000s and therefore we compared it against simulated AGB for the years 2009-2011. We regridded all data set to a $0.5^{\circ} \times 0.5^{\circ}$ resolution. In addition, we used maps of PFTs as derived from the ESA CCI land cover map V2.0.7 (Li et al., 2018; Forkel et al., 2014).

2.6 Evaluation metrics

To quantify the performance of the model output, we applied the Pearson Correlation between two time series, the normalized mean square error (NMSE; Kelley et al., 2013) and the Willmott coefficient of agreement (W; Willmott, 1982) to describe differences between the model simulation and the reference data sets. The NMSE is calculated by:

$$NMSE = \frac{\sum_{i=1}^{N} (y_i - x_i)^2}{\sum_{i=1}^{N} (x_i - \bar{x})^2}$$
(15)

where y_i is the simulated and x_i the observed value in the grid cell i. \bar{x} is the mean observed value. The NMSE is zero for perfect agreement between simulated and modelled results, 1.0 if the model is as good as using the observed mean as a predictor and larger than 1.0 if the model performs worse than that. We chose the NMSE to represent and compare the model errors, as it has a squared error term, which puts a stronger emphasis on large deviations between simulations and observations as compared to a linear term, and due to its normalization it is comparable across different parameters. Especially for fire simulations we have a relatively large deviation between simulations and observations.

15 The Willmott coefficient of agreement is given by:

$$W = 1 - \frac{\sum_{i=1}^{N} (y_i - x_i)^2 \cdot A_i}{\sum_{i=1}^{N} (|y_i - \bar{x}| + |x_i - \bar{x}|)^2 \cdot A_i}$$
(16)

which additionally accounts for the area weight A_i of the grid cell i. The Willmott coefficient is a squared index, where a value of 1 stands for perfect agreement between simulated and modelled runs and gets smaller for worse agreements with a minimum of 0. Unlike the coefficient of determination, the Willmott coefficient is additionally sensitive to biases between simulations and observations.

3 Results

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3.1 Performance of optimized fire danger index formulations

Overall, the yearly burnt area simulated by the standard SPITFIRE model (using the original Nesterov index, NI_{orig}) showed poor simulation results over South America as compared to the GFED4 evaluation data set (Fig. 4 a and b: NMSE=1.80, W=0.27). The average yearly burnt area (without croplands) for South America was with ca. 14 million ha about 25 % smaller than the observed burnt area with 19 million ha in the shown period from 2005-2015. The spatial pattern of the modelled burnt area agreed well with the GFED4 data in the region of the Cerrado that is close to the Caatinga border, while the fires in other semi-arid regions of the continent were underestimated. For example, simulated fire is underestimated in the savanna-areas in the northern part of South America (on the Columbian-Venezuelian border) even though there is a strong signal visible in

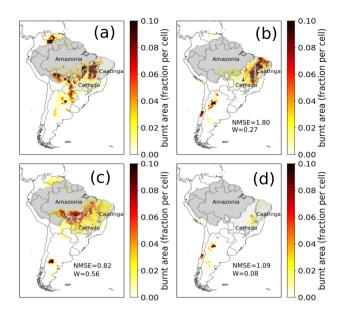


Figure 4. Yearly burnt area over a mean from 2005-2015 as fraction per cell. (a) GFED4 evaluation data of burnt area excluding crops and simulated burnt area by SPTIFIRE using the (b) NI_{orig} version, (c) VPD_{optim} version, (d) NI_{optim} version

the satellite observations. The biomes Caatinga and Cerrado, which are of special interest in this study, showed very different results: while fire in Caatinga was overestimated, it was underestimated in the Cerrado.

The optimized version using NI_{optim} (Fig. 4 d), led to an overall decrease of fire, with a slight improvement of NMSE (1.09) as compared to NI_{orig} and a worse Willmott coefficient of 0.08. While the overestimation of fire in Caatinga was reduced, all the fires across South America also have decreased significantly, which led to a general underestimation of fire by 90 % (2 million ha). The optimized version, using VPD_{optim} (Fig. 4 c), clearly improved the model performance, mainly by shifting much of the simulated burnt area from the sparsely vegetated Caatinga towards the Cerrado region (NMSE=0.82 and W=0.56). In addition, by using VPD_{optim} , the model results also showed fire occurrence in northern South America, where fire was not at all or only minimally simulated when using NI_{optim} or NI_{orig} . The total burnt area was in this model version ca 20 % smaller than the evaluation data set (16 million ha).

The burnt area time series from 2005 to 2015 provides a more detailed view on the model performance for the fire-prone Cerrado and Caatinga region (Fig. 5). While model performance was relatively good for the Cerrado region with NI_{orig} (NMSE=0.3, W=0.89, R^2 =0.78), the simulated burnt area was strongly overestimated in the Caatinga region throughout the whole period (NMSE=327.82, W=0.14, R^2 =0.59). After the optimization of the NI, the model performance indeed improved for the Caatinga (NMSE=1.07, W=0.73, R^2 =0.31), but at the same time the performance for the Cerrado declined (NMSE=1.07, W=0.36, R^2 =0.4). On the other hand VPD_{optim} showed an improved fire representation compared to the standard settings in the Cerrado (NMSE=0.27, W=0.9, R^2 =0.8) as well as in the Caatinga (NMSE=15.2, W=0.46, R^2 =0.56). Even though fire in the Caatinga was still overestimated, the NMSE decreased by a factor of six.

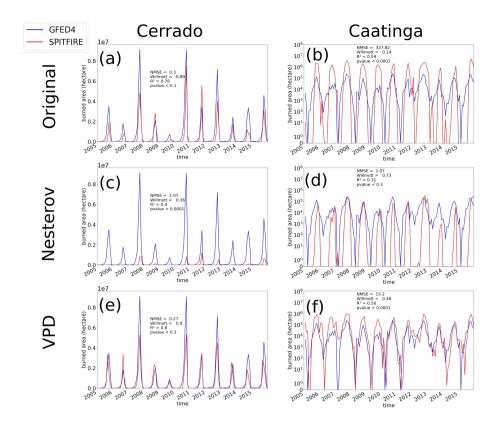


Figure 5. Time-series of monthly burnt area from 2005 - 2015 simulated by SPITFIRE (red lines) compared to GFED4 evaluation data (blue lines) for: (a) The Cerrado region, using NI_{orig} . (b) The Caatinga region, using the NI_{orig} . (c) The Cerrado region, using NI_{optim} . (d) The Caatinga region, using NI_{optim} . (e) The Cerrado region, using VPD_{optim} . (f) The Caatinga region, using VPD_{optim} . Note the logarithmic scale for the Caatinga, which was applied in order to account for the large differences between the different model versions (for a non-logarithmic version see Fig. A6).

Overall, the total amount of burnt area in the Cerrado was for all three model versions smaller than in the evaluation data set. Fire occurrence in the Caatinga was, on the other hand, largely overestimated by the NI_{orig} and the VPD_{optim} version. Just in the NI_{optim} version the burnt area of the Caatinga is in the same order of magnitude as the evaluation data set, which also led, however, to a large underestimation in the Cerrado (Tab. 1 and Fig. 4). Also the Amazonia region mostly improved by using the VPD_{optim} version (Tab. 1, Fig. A3). The R^2 and the Willmott coefficient improved, while the NMSE increased slightly. With the Nesterov index fire was strongly underestimated in the Amazon region, while the optimized VPD fixes this underestimation. The fire is only modelled (and also observed, see Fig. 4) at the edges to the Amazon, where wood density is lower and deforestation takes place. In the closed continuous forest area towards the center of the Amazon almost no fire is observed and neither modelled. The total burnt area increased from 0.7 million ha to 4.8 million ha (for VPD_{optim}), which is now a bit overestimated to the observed burnt area of 3.4 million ha. Using the NI_{optim} all error metrics as well as the total burnt area decreased.

Table 1. Comparison of the burnt area results in terms of NMSE, the Willmott coefficient of agreement and the sum (in ha per year) between NI_{orig} , VPD_{optim} , NI_{optim} and the GFED evaluation data

Region	NMSE	Willmott	Sum
Spatial - South America			
GFED			$1.9\cdot 10^7$
NI_{orig}	1.80	0.27	$1.4\cdot 10^7$
VPD_{optim}	0.82	0.56	$1.6\cdot 10^7$
${ m NI}_{optim}$	1.09	0.08	$0.2 \cdot 10^7$
Temporal - Cerrado			
GFED			$9.2\cdot 10^6$
NI_{orig}	0.30	0.89	$5.2\cdot 10^6$
VPD_{optim}	0.27	0.90	$6.4\cdot10^6$
NI_{optim}	1.07	0.36	$0.6\cdot10^6$
Temporal - Caatinga			
GFED			$0.4\cdot 10^6$
NI_{orig}	327.82	0.14	$6.0\cdot 10^6$
VPD_{optim}	15.2	0.46	$1.6\cdot 10^6$
NI_{optim}	1.07	0.73	$0.3 \cdot 10^6$
Temporal - Amazonia			
GFED			$3.4\cdot 10^6$
NI_{orig}	0.83	0.56	$0.7\cdot 10^6$
VPD_{optim}	0.93	0.83	$4.8\cdot 10^6$
NI_{optim}	1.22	0.32	$0.02\cdot10^6$

3.2 Optimized model parameters

Seven fire-related parameters were optimized, in order to improve the fire representation in the LPJml4-SPITIFRE model. Here we compare the optimized parameters for the different model versions in order to evaluate and discuss parameter variability and changes. Table 2 shows all parameters, used for the optimization, their lower and upper boundary and the resulting optimized value. Since the FDI directly controls the amount of modelled fire, the FDI scaling factors for the different PFTs are central for this analysis. For both optimization experiments the boundaries were, hence, set rather generously within one magnitude of the original value. In the NI_{optim} experiment, all scaling factors generally decreased compared to the standard values used for NI_{orig} . Here, TrH displayed the smallest scaling factor (9.39 · 10⁻⁶), followed by TrBE (2.48 · 10⁻⁵) and TrBR (4.76 · 10⁻⁵). Since the VPD is a newly implemented fire danger index, we have no standard values to compare the optimized scaling factors with. Here, the TrBE showed the largest value (22.41), ca. 20 times as large as the TrBR (1.21) and TrH (1.13) (Tab. 2).

Table 2. All optimized parameters with their standard values, the upper and lower boundary of the parameter ranges and the resulting optimized value including parameter for specific PFTs and global parameter, which have the same value for all PFTs. All parameters exept ρ_b have no unit.

Parameter	PFT	Standard value (as in	Lower	Upper	After optimiza-
		Thonicke et al., 2010)	bound.	bound.	tion
\mathbf{NI}_{optim}					
scaling factor FDI α_{NI_i}	TrBE	$3.34 \cdot 10^{-5}$	$7 \cdot 10^{-6}$	$1.33\cdot10^{-4}$	$2.4885 \cdot 10^{-5}$
scaling factor FDI α_{NI_i}	TrBR	$3.34 \cdot 10^{-5}$	$7\cdot 10^{-6}$	$1.33\cdot 10^{-4}$	$4.7649 \cdot 10^{-5}$
scaling factor FDI α_{NI_i}	TrH	$6.67 \cdot 10^{-5}$	$7\cdot 10^{-6}$	$1.33\cdot 10^{-4}$	$9.3949 \cdot 10^{-6}$
fire duration parameter p_t	all PFTs	-11.06	-13	-9	-9.0011
scorch height parameter F	TrBE	0.1487	0.01	0.6	0.1282
scorch height parameter F	TrBR	0.061	0.01	0.6	0.0752
crown mortality parameter rCK	TrBE	1.0	0.5	1	0.5030
crown mortality parameter rCK	TrBR	0.05	0	0.5	0.4038
fuel bulk density ρ_b (kg m ⁻³)	TrBE	25.0	22.5	27.5	26.6473
fuel bulk density ρ_b (kg m ⁻³)	TrBR	13.0	11.7	14.3	13.1896
fuel bulk density ρ_b (kg m ⁻³)	TrH	2.0	1.8	2.2	2.0019
human ignition parameter p_h	all PFTs	-0.5	-0.6	-0.4	-0.5426
surface intensity threshold	all PFTs	10^{-6}	10^{-7}	10^{-5}	$1.0317 \cdot 10^{-6}$
VPD _{optim}					
scaling factor FDI α_{VPD_i}	TrBE	-	1	50	22.4181
scaling factor FDI α_{VPD_i}	TrBR	-	1	50	1.2135
scaling factor FDI α_{VPD_i}	TrH	-	1	50	1.1299
fire duration parameter p_t	all PFTs	-11.06	-13	-9	-11.3753
scorch height parameter F	TrBE	0.1487	0.01	0.6	0.1930
scorch height parameter F	TrBR	0.061	0.01	0.6	0.0799
crown mortality parameter rCK	TrBE	1.0	0.5	1	0.9983
crown mortality parameter rCK	TrBR	0.05	0	0.5	0.4801
fuel bulk density ρ_b (kg m ⁻³)	TrBE	25.0	22.5	27.5	22.5923
fuel bulk density ρ_b (kg m ⁻³)	TrBR	13.0	11.7	14.3	13.3750
fuel bulk density ρ_b (kg m ⁻³)	TrH	2.0	1.8	2.2	1.8944
human ignition parameter p_h	all PFTs	-0.5	-0.6	-0.4	-0.5332
surface intensity threshold	all PFTs	10^{-6}	10^{-7}	10^{-5}	$3.6317 \cdot 10^{-6}$

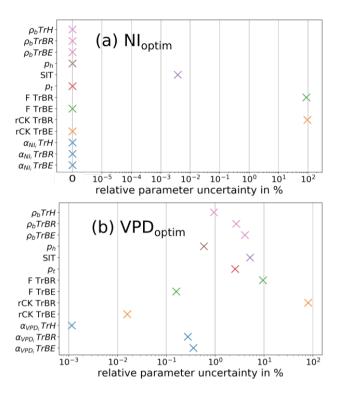


Figure 6. Relative uncertainty of model parameters after optimization for (a) NI_{optim} and (b) VPD_{optim}. The relative uncertainty is the ratio of the uncertainty after the optimization (range of all parameter sets with low cost, below the 0.05 quantile) divided by the uncertainty before the optimization (range of the parameters for the optimization). Low and high values of relative uncertainty indicate strongly and weakly constrained parameters, respectively. SIT denotes the surface intensity threshold. PFT-dependent parameters are grouped with the same color.

In case of the other optimized parameters the boundaries were set smaller in order to decrease the possibility that a large error in the estimation of several parameters would lead to a better overall cost in the optimization procedure. The human ignition parameter became smaller for both optimizations, which led to a smaller amount of human ignitions (from -0.5 to -0.54 in NI_{optim} and -0.53 in VPD_{optim}). The fuel bulk density increased for all three tropical PFTs in the NI_{optim} version, while for VPD_{optim} the fuel bulk density of the TrBE and TrH PFTs decreased and for the TrBR increased. For the NI_{optim} version, the fire duration parameter (p_t) increased, leading to a shorter fire duration (from -11.06 to -9), while the value for the VPD_{optim} version stayed relatively similar (-11.37) to the prior value. The surface intensity threshold became slightly larger for the NI_{optim} version than the original value (from 10^{-6} to $1.03 \cdot 10^{-6}$). For VPD_{optim} the parameter increased by a factor of three (3.63 · 10^{-6}). The mortality related parameters F and FCK led in the FICK version both to a decrease in the fire-related mortality for FICK and an increase for FICK which led to an increased mortality.

The relative uncertainties were for most optimized parameters very small (between 0 and 10%), hence these parameters were strongly constrained (Fig. 6). Just the fire-mortality related parameters (F and rCK) had large uncertainties for the TrBR, hence

Table 3. Comparison of the results for AGB and the TrBE PFT cover in terms of NMSE and the Willmott coefficient of agreement between NI_{orig} , VPD_{optim} and NI_{optim} in South America (SA), in the Cerrado and in the Caatinga.

Region	NMSE	Willmott
AGB		
$\mathrm{SA}\left(\mathrm{NI}_{orig} ight)$	0.97	0.83
$SA (VPD_{optim})$	0.91	0.84
$SA\left(NI_{optim} ight)$	0.99	0.83
Cerrado (NI_{orig})	15.06	0.25
Cerrado (VPD_{optim})	12.36	0.28
Cerrado (NI_{optim})	16.06	0.24
Caatinga (NI_{orig})	11.93	0.32
Caatinga (VPD_{optim})	8.57	0.36
Caatinga (NI_{optim})	10.44	0.33
FPC - Evergreen (TrBE)		
$\mathrm{SA}\left(\mathrm{NI}_{orig} ight)$	0.42	0.82
$SA (VPD_{optim})$	0.41	0.82
${ m SA}\left({ m NI}_{optim} ight)$	0.43	0.81
Cerrado (NI_{orig})	1.04	0.60
Cerrado (VPD_{optim})	0.70	0.64
Cerrado (NI_{optim})	1.40	0.55
Caatinga (NI_{orig})	1.73	0.40
Caatinga (VPD_{optim})	1.54	0.29
Caatinga (NI_{optim})	2.05	0.44

were weakly constrained. For VPD_{optim} the uncertainty of rCK (TrBR) was 0.8 and for NI_{optim} the uncertainty of F was 0.9 and for rCK 1 (TrBR).

The decrease in the model error (cost) due to the optimization process has been mainly due to improvement in the burnt area. While for the NI_{optim} the cost of the burnt area dataset improved by 81%, the cost of the biomass dataset improved just by 6%. In case of the VPD_{optim} the cost of the burnt area dataset improved by 49%, whereas the biomass dataset improved by 19% (Fig. A5).

3.3 Model evaluation for South America

The modelled above-ground biomass (AGB) of trees in South America was throughout all model versions larger than the evaluation data set indicates (Fig. 7). Especially the biomass in the Amazon region is with an average of ca. 20 kgC/m² about one third overestimated. The drier savanna regions on the continent yielded a biomass of ca. 5-10 kgC/m², which also

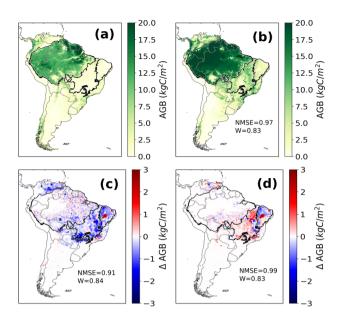


Figure 7. Annual above ground biomass (AGB) of trees over a mean from 2005-2015 in kgC/m². (a) Avitabile evaluation data. (b) Simulated AGB by LPJmL4-SPITFIRE in the NI_{orig} version. (c) Difference between VPD_{optim} and NI_{orig}. (d) Difference between NI_{optim} and NI_{orig}. Red (blue) color indicates a larger (smaller) biomass after the optimization.

constitutes an overestimation in wide parts of the Cerrado and the Caatinga biome (evaluation data shows between 1-5 kgC/m², also see Roitman et al., 2018).

The differences among the different model versions are marginal: The VPD $_{optim}$ version had the best performance compared to the evaluation data set (NMSE=0.91, W=0.84), the NI $_{orig}$ version had the second best performance (NMSE=0.97, W=0.84) and the NI $_{optim}$ the worst performance (NMSE=0.99, W=0.83). The model optimization scheme focuses on fire parameters, hence the model performance for AGB can only improve in areas, where the fire occurrence has been modelled poorly and the vegetation-fire interactions have improved due to the optimization process. For example in the center of the Amazon rainforest almost no fire is found in the evaluation data nor is simulated. Hence no improvement of burnt area as well as AGB can be achieved. On the other hand, in regions where the modelling error of burnt area is now reduced, this can also improve simulated AGB, hence vegetation-fire interactions. In the fire-prone Caatinga and Cerrado the VPD $_{optim}$ version mostly decreased the biomass by up to 3 kgC/m², showing a better performance compared to the evaluation data set (e.g. in the Cerrado the NMSE decreased from from 15.06 to 12.36 in the VPD $_{optim}$ version compared to NI $_{orig}$, see Tab. 3).

The modelled foliage projective cover (FPC) showed for all three model versions a strong underestimation compared to the evaluation data set of the TrBE throughout the whole Amazonian region (ca. 50% compared to ca. 100% in the evaluation dateset). In the fire-prone biomes Cerrado and Caatinga, however, the TrBE PFT was sometimes overestimated (TrBE cover between 0 and 40 %, Fig. 8). In the regions with less TrBE the dominant PFT was mostly TrBR (Cerrado) or TrH (Caatinga) (see Fig. A1 and A2).

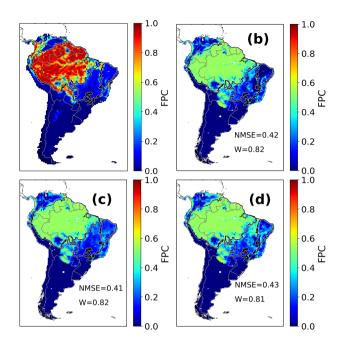


Figure 8. Annual FPC cover by tropical broadleaved evergreen PFT over a mean from 2005-2015 as fraction per cell. (a) ESA-CCI evaluation data (b) Simulated FPC by LPJmL4-SPITFIRE using the NI_{orig} version (c) Simulated FPC by LPJmL4-SPITFIRE using the VPD_{optim} version (d) Simulated FPC by LPJmL4-SPITFIRE using the NI_{optim} version

 NI_{optim} led to an overall decrease in the model performance also in terms of the TrBE distribution, as both, the NMSE and the Willmott coefficient declined compared to NI_{orig} (NI_{orig} : NMSE=0.42, W=0.82; NI_{optim} : NMSE=0.43, W=0.81).

The VPD_{optim} version, on the other hand, showed an slightly improved TrBE distribution (NMSE=0.41, W=0.82) but also in this case we obtained an even larger improvement, when only the fire-prone regions Cerrado or Caatinga are considered (Tab. 3). Also for the TrBR and TrH PFT distributions the optimization lead to an improved performance using the VPD_{optim} in the Caatinga and Cerrado, whereas the PFT distribution in the Amazon remained similar to the prior PFT distribution. In the NI_{optim} version, parameter optimization only slightly reduced TrBR cover showing a worse performance compared to VPD_{optim}. However, herbaceous cover changed only slightly in all optimization experiments (Fig. A1 and A2).

4 Discussion

In summary, our results show that the implementation of a new fire danger index based on the water vapor pressure deficit FDI_{VPD} and its optimization against satellite data sets improved the simulations of fire in LPJmL4-SPITFIRE, both in terms of spatial patterns as well as temporal dynamics of burnt area. In the following, we discuss the model improvements, limitations and recommendations for future improvements of process-based global fire models within the DGVM framework.

4.1 Improvements in model performance

The VPD results showed a better model performance for fire in the spatial dimension, as well as in the temporal dimension (Tab. 1 and 3). Compared to the Nesterov index, FDI_{VPD} uses additional climate input as relative humidity and precipitation. In the calculation of the Nesterov index precipitation is just used as a threshold. This leads to a better accounting of the very different climatic conditions among various biomes. Furthermore, the FDI_{VPD} includes a direct representation of the vegetation density. The significance of this has been recently shown by findings of Forkel et al. (2019a) who have emphasized the importance of past plant productivity and fuel production for burnt area. This is particularly important for differentiating between fires in biomes with similar PFT distribution. For example, the vegetation density is much larger in the Cerrado, even though the Caatinga and Cerrado have a similar modelled PFT composition, which provides more fuel and therefore leads to a higher fire danger.

While the seasonal and interannual variability in the Caatinga has improved largely using the FDI_{VPD} (NMSE decreased by a factor of ca. 20), the improvement in the Cerrado was relatively small (NMSE decreased by ca. 10%). This is due to the fact, that the optimization tries to obtain a compromise between the different optimized cells. As the model performance was originally much better for the Cerrado, the largest improvement could be achieved for the Caatinga. We have also chosen a large amount of cells in the Caatinga, because the model performance was here particularly bad. This leads to a large improvement in the time series of the Caatinga region, while the improvement for the Cerrado was less significant. With the Nesterov index fire was strongly underestimated in the Amazonia region, while the optimized VPD increases the modelled burnt area. The fire is only present at the edges of the Amazon (both in model and observation, see Fig. 4), where tree density is lower and deforestation takes place. In the closed continuous forest area towards the center of the Amazon almost no fire is observed and also not simulated.

Another result of the optimizing procedure, using FDI_{VPD} , was the improvement of the PFT distribution and the above-ground biomass of trees especially in the fire-prone biomes Caatinga and Cerrado (Fig. 8). For example, the central Amazon, where fire is a scarce event, shows almost no changes compared to the non-optimized model version. Here, it is the improvement of the vegetation model itself, and not the fire module, which can help to improve the model performance of LPJmL4-SPITFIRE. Hence, it emphasizes that we need to include further parameters in the optimization, which impact directly the PFT distribution, biomass and fire to obtain a significant improvement in the spatial and temporal distribution of both, vegetation and fire. However, this study focused solely on the parameters within the SPITFIRE module. Due to the focus on fire related parameter, the cost of the burnt area dataset decreases much more than the cost of the biomass dataset (Fig. A5). Hence we only get a substantial improvement in model performance in semi-arid, fire-prone biomes, where vegetation dynamics and fire are strongly coupled.

During the optimization-process most of the optimized parameters were well constrained, except for the mortality-related parameters for the TrBR PFT (Fig. 6). The TrBR PFT is dominant in the fire-prone regions, where the mortality-related parameters have a large impact on vegetation dynamics. Hence, they impact multiple LPJmL routines, which are responsible for the PFT distribution and carbon cycling. This leads in turn to a less certain parameter estimation. In order to better constrain

these parameters also the optimization of vegetation model parameters would be necessary to decrease the uncertainties. The fire danger index scaling factors (α_{NI_i} and α_{VPD_i}) convert the quantified fire risk (NI or VPD) into the actual fire danger (FDI). Both scaling factors thus set the magnitude of the fire danger for the different PFTs. Hence they impact directly the fire spread, burnt area and the number of fires as well as indirectly fire mortality. These very important parameters vary significantly

for the different PFTs. TrH has the smallest scaling factor in case of both FDIs, which leads to a lower fire danger compared to the other PFTs. This indicates a prior overestimation of the fire danger of grass in tropical South America, as grasslands are generally parametrized to have a low fire resistance and moisture content and can hence burn very easily. This overestimation, compared to tree PFTs has been decreased by the optimization. In case of the VPD also the TrBR is scaled by a much smaller factor than the TrBE, which leads to a lower fire danger index. This is due to the fact, that the TrBR is dominant in dry and fire-prone regions, which experience frequent fires. Here the burnt area was often overestimated by SPITFIRE (e.g. Caatinga or eastern Cerrado) and is now decreased. On the other hand, a larger FDI for the TrBE allows more fire in wetter regions at the edge between the Cerrado and the Amazon rainforests, where TrBE is more dominant. The mortality risk of TrBE for VPD_{optim} remains close to the prior value of 1, confirming previous assumptions about its fire sensitivity. Whereas the rCK for TrBR increased to 0.48, close to the upper boundary, meaning that a mortality risk of 50% when the full crown is scorched and a 7% mortality risk when 50% of the crown is scorched, which makes the TrBR less resistant against crown damage than before. Due to this changes the overestimation of biomass in the original model for the Cerrado/Caatinga region decreased (see Fig. 7).

4.2 Limitations of the optimization process

Generally, optimizing a model against burned area is challenging because 1) of the skewed statistical distribution of burned 20 area and 2) because temporal or spatial mismatches in simulated burning can cause large model-data errors. These issues can be avoided with the choice of an appropriate cost function. For example, squared-error metrics tend to underestimate the variance of burned area in comparison to, e.g., the Kling-Gupta efficiency as it has been shown in the optimization of an empirical model for burned area (see Tab. A3 in Forkel et al. (2017)). Here, the optimum parameter set for the Nesterov index-based model resulted in almost no fires across South America. Thereby the optimization algorithm tries to decrease the model error by tending towards a conservative 'no fire strategy' for all biomes. This result nicely demonstrates the need to evaluate model optimization results against spatially and temporally independent data and independent variables (Keenan et al., 2011).

The Nesterov index is not able to capture fire variability within the Caatinga as well as the Cerrado at the same time. This shows that the difference in the PFT distribution between these two biomes is not adequately modelled by LPJmL or just using PFT dependent scaling factors did not sufficiently improve the model performance when using the Nesterov index. On the other hand, using the VPD fire danger index reduced the model error for burned area in both biomes, by improving the modelled performance for the Caatinga and maintaining the good performance of the Cerrado region. Since improved performance of the fire model mainly had minor effect on improving FPC of the tropical PFTs, the presented optimization scheme has to go along with process-based improvements in both, in the fire and in the vegetation modules of LPJmL.

Fire largely depends on the vegetation type and their associated flammability, fire tolerance and mortality. Hence an accurately

modelled vegetation distribution is crucial for a good model performance in terms of burnt area and fire effects (Forkel et al., 2019a; Rogers et al., 2015). As shown in Fig. 8, A1 and A2, the modelled PFT coverage showed an equal distribution of tropical raingreen and evergreen PFTs throughout wide parts of central-northern South America. Evaluation data shows, however, an TrBE dominance in the wet rainforest regions and a TrBR dominance in the Cerrado and Caatinga. This emphasizes the potential to improve the fire modelling further, based on an improved PFT distribution. In the tropical rainforest the TrBR proportion is overestimated, which leads to problems in the optimization procedure, since TrBR has very different effects on fire spread and is more fire-tolerant (different fuel characteristics and resulting fire intensity). This leads to a lower fire-related mortality, which fits better to the drier and fire prone savanna-like regions (e.g. Cerrado). The poorly modelled PFT distribution also is responsible for the overestimation of the burnt area in the Amazon region. Because of the too large fraction of TrBE in the Cerrado/Caatinga region the scaling factor for this PFT is relatively high. This leads in turn to an overestimation in the Amazon region, where the fraction of the TrBE is larger.

Since the offset is very small, the years 2000-2003 (first three years of GLDAS 2.1, before the optimization period) are enough for the model to recover from the offset and the carbon pools to return to equilibrium. To exclude the possibility that long-term trends within GLDAS 2.0 changed the modelled vegetation state significantly, we tested our optimization also just based on GLDAS 2.0 data (until 2010) and on GLDAS 2.1 data (2000-2017) only, using the same years for model spinup, optimization and evaluation. Both versions yielded similar results compared to the optimization presented in this study (results not shown). Due to the fact that evaluation data are only available for the last 10-20 years, we are constrained to optimize the model in this relatively short time period. In South America these years were subject to an unusual high amount of severe droughts and other extreme events (Panisset et al., 2017). As a result, an optimization in this period could lead to a worse model performance in a period with less pronounced droughts. This is due to the non-linear relationship between the drought signal in the input data set and the resulting modelled biosphere behavior. Nonetheless, we were able to improve the interannual variability and hence, the model performance to a great extent for the Caatinga and slightly for the Cerrado and Amazon regions (Fig. 5 and A3). The Cerrado already had a very good modelling performance before the optimization process, which now only slightly improved. The performance of the interannual and seasonal variability of burnt area for total South America improved substantially (Fig. A3). The optimized SPITFIRE is now better able to simulate accurately the climate dependent seasonal and interannual variability and the climate depend

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Systematic optimizations within a model-data integration setup of fire models which are embedded in a DGVM are still very rare. Previously, Rabin et al. (2018) optimized the fire model FINAL.1 within the land-surface model LM3. Our study differs from Rabin et al. (2018) in the conceptual design of the vegetation-fire models and the optimization process. While LM3 has been run on a 2° longitude by 2.5° latitude, which is much coarser than LPJmL with 0.5° by 0.5°. This difference allows us to account for a locally better climate input, vegetation and fire interaction. While FINAL.1 is a process-based model, many calculations (e.g. the fire spread routine) are done by multiplying the important factors and fitting the resulting values to observational data. SPITFIRE tries to model the important fire variables by simulating the underlying processes, and by taking the influence of climate and the different fire ignitions into account. An advantage of FINAL.1 is the inclusion of agricultural fires based on a statistical approach. Whereas Rabin et al. (2018) used a local search algorithm (Levenberg-Marquardt algorithm) to

ability as well as the spatial extent of fire on natural land throughout the fire-prone woodlands of South America.

optimize their model, we used a global search algorithm (genetic optimization). Local search algorithms depend on the chosen initial parameter sets and might eventually end up in a local optimum. A genetic optimization algorithm allows to explore the full parameter space and hence gives a higher chance to find the global optimum. However, local search algorithms require less iterations than global search algorithms (300 in Rabin et al. (2018) vs. 16000 in our study). Forkel et al. (2014) tested the optimization of LPJmL with different optimization algorithms and found that it was not feasible to optimize LPJmL with a local search algorithm. Rabin et al. (2018) ran the model during the optimization process only for the period of 1991-2009, whereas we made complete model runs including 5000 years of spinup in order to get a model equilibrium for each tested parameter combination.

4.3 Limitations of fire modelling in LPJmL4-SPITFIRE

- In fire-prone regions the interactions between fire and vegetation dynamics are strong, hence are posing a challenge for global fire models embedded in DGVMs. By just focussing on fire-related parameters, an optimization approach can only to a certain extent improve PFT distribution and simulated biomass. For a good fire representation e.g. in the Cerrado and Caatinga, a shrub PFT could further improve the model performance. Most fires in this region occur, where shrub PFTs are abundant. LPJmL tries to account for this by establishing rather small raingreen PFTs as a shrub replacement. A much better option would be a separate shrub PFT with parameters leading to a high flammability, but also a low fire mortality. An optimization of LPJmL4-SPITFIRE, including shrub PFTs could yield better results than shown in this study.
 - Fire models embedded in DGVMs should build on a FDI which is complex enough to account for various fire dynamics, while it's parameterization should be simple enough to be accurately applied on a global scale. While the VPD is more complex and takes into account more climatic input as the Nesterov index, it is still relatively easy to implement in a global fire model.
- There are various other fire danger indices used for modelling purposes, as well as real fire danger assessment and fire forecast purposes. For example, fire-prone countries have developed their own fire danger indices (e.g. Canada, Australia), which are suited to the unique local fire regimes and vegetation dynamics. In a global modelling approach, however, we need to find one fire danger index, which suits best for all regions of the world and has a relatively easy implementation to decrease computational cost and the number of input data sets (which might be unavailable or uncertain).
- Currently, SPITFIRE does not account for fire in managed land like cropland or managed grassland. We accounted for this by excluding cropland fires from the evaluation burnt area data set. We do, however, not account for the proportion of grassland, which is used for e.g. cattle ranching. Since in SPITFIRE fire is not enabled on pasture, our results show a slightly smaller burnt area throughout South America than could be expected with managed land included and hence also compared to the GFED4 evaluation data set. This effect is however small, because pasture lands cover a substantial fraction only in very few grid cells (e.g. southern Cerrado; Parente et al., 2017). Fire on managed land is generally difficult to predict in a DGVM because the reason and timing of using fire depends less on climatic factors but mostly on social and political decisions which can vary between countries, regions and localities. We expect further improvement of model performance especially in regions of large land-use areas with fires on pastures included (e.g. Rabin et al., 2018; Pfeiffer et al., 2013).

5 Conclusions

We significantly improved the fire representation within LPJmL4-SPITFIRE, applied for South America, by implementing a new fire danger index and applying a model-data integration setup to optimize fire-related parameters. We improved the seasonal and interannual variability, as well as the spatial pattern of burnt area in South America. In addition, modelling of related vegetation variables, e.g., the biomass and the PFT distribution in the fire-prone Cerrado and Caatinga biomes have also been improved.

Optimizing fire parameters has its limits due to error propagation of the PFT distribution and hence their fire traits influencing simulated fire spread and behavior. Furthermore, it remains a challenge to find a fire danger index that is physically interpretable and can be applied globally. In this study, the parameter-optimization by using FDI_{NI} led to a large underestimation of fire and a generally worse model performance, when focusing on the Cerrado and Caatinga biome. However, implementing the more complex FDI_{VPD} and optimizing it thereafter, led to an improved model performance compared to the original SPITFIRE implementation for South America. Our results demonstrate that the improvement of model processes, as well as a systematic model-data optimization are required in order to obtain a more accurate fire representation within complex DGVMs, where observations or experimental evidence to constraint fire parameter are scarce. This work highlights the potential for future model-data integration approaches to obtain a better fire model performance in a global setting, based on improved vegetation dynamics within LPJmL4.

Code availability. Upon request. With publication of this article the model code will be published at https://github.com/PIK-LPJmL, analogously to Schaphoff et al. (2018a). We will publish the code of LPJmLmdi along with the model code on the github page of LPJmL: https://github.com/PIK-LPJmL/LPJmLmdi

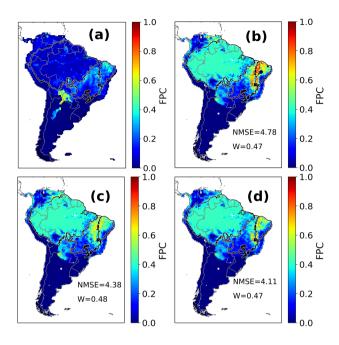


Figure A1. Annual FPC cover by tropical broadleaved raingreen PFT over a mean from 2005-2015 as fraction per cell. (a) ESA-CCI evaluation data (b) Simulated FPC by LPJmL4-SPITFIRE using the NI_{orig} version (c) Simulated FPC by LPJmL4-SPITFIRE using the VPD_{optim} version (d) Simulated FPC by LPJmL4-SPITFIRE using the NI_{optim} version

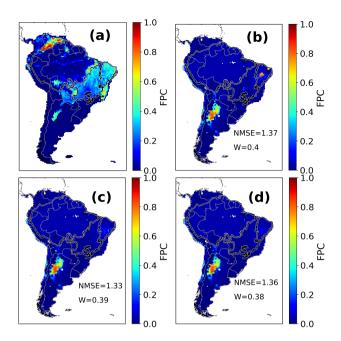


Figure A2. Annual FPC cover by tropical herbaceous PFT over a mean from 2005-2015 as fraction per cell. (a) ESA-CCI evaluation data (b) Simulated FPC by LPJmL4-SPITFIRE using the NI_{orig} version (c) Simulated FPC by LPJmL4-SPITFIRE using the VPD_{optim} version (d) Simulated FPC by LPJmL4-SPITFIRE using the NI_{optim} version

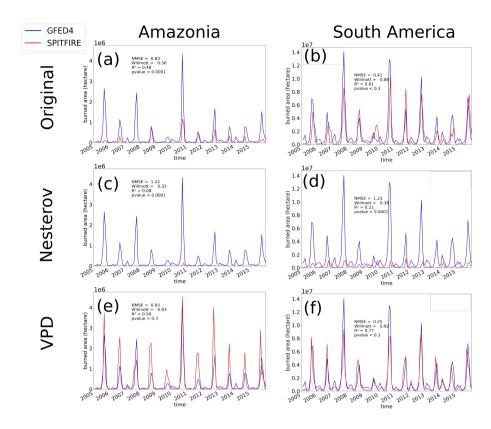


Figure A3. Time-series of monthly burnt area from 2005 - 2015 simulated by SPITFIRE (red lines) compared to GFED4 evaluation data (blue lines) for: (a) The Amazonia region, using NI_{orig}. (b) Total South America, using the NI_{orig}. (c) The Amazonia region, using NI_{optim}. (d) Total South America, using NI_{optim}. (e) The Amazonia region, using VPD_{optim}. (f) Total South America, using VPD_{optim}.

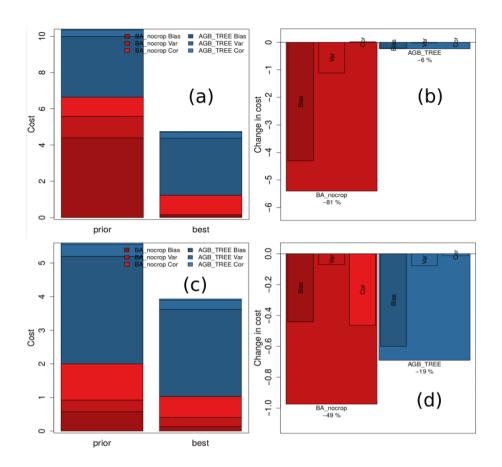


Figure A4. Cost reduction of the burnt area and the biomass during the optimization process, by showing the various components of the cost that are related to model-data bias, variance ratio and correlation. The cost for burnt area for NI_{optim} decreased by ca. 81%, whereas the cost of the biomass only decreases by ca. 6% (a and b). For VPD_{optim} the cost decreased by ca. 48% for burnt area and ca. 19% for the biomass (c and d). Hence the impact of the optimization process on burnt area is much larger due to the focus on fire parameters.

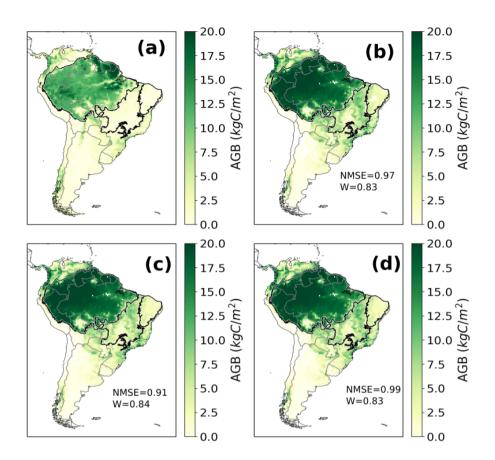


Figure A5. Annual above ground biomass (AGB) of trees over a mean from 2005-2015 in kgC/m². (a) Avitabile evaluation data. (b) Simulated AGB by LPJmL4-SPITFIRE in the NI_{orig} version. (c) Simulated AGB by LPJmL4-SPITFIRE in the VPD_{optim} version. (d) Simulated AGB by LPJmL4-SPITFIRE in the NI_{optim} version.

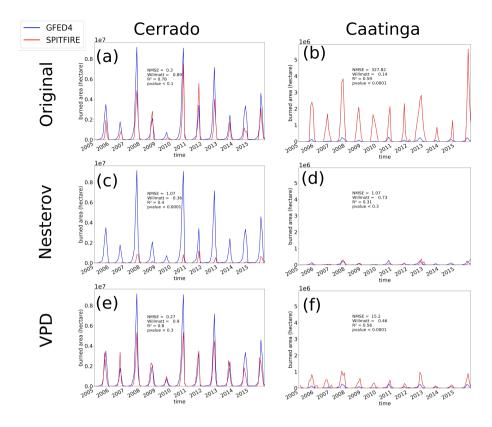


Figure A6. Time-series of monthly burnt area from 2005 - 2015 simulated by SPITFIRE (red lines) compared to GFED4 evaluation data (blue lines) for: (a) The Cerrado region, using NI_{orig} . (b) The Caatinga region, using the NI_{orig} . (c) The Cerrado region, using NI_{optim} . (d) The Caatinga region, using NI_{optim} . (e) The Cerrado region, using VPD_{optim} .

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Competing interests. The authors declare that they have no conflict of interest.

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References

- Alvares, C. A., Stape, J. L., Sentelhas, P. C., de Moraes Gonçalves, J. L., and Sparovek, G.: Köppen's climate classification map for Brazil, Meteorol. Z., 22, 711–728, https://doi.org/10.1127/0941-2948/2013/0507, 2013.
- Andela, N., Morton, D. C., Giglio, L., Chen, Y., van der Werf, G. R., Kasibhatla, P. S., DeFries, R. S., Collatz, G. J., Hantson, S., Kloster, S.,
 Bachelet, D., Forrest, M., Lasslop, G., Li, F., Mangeon, S., Melton, J. R., Yue, C., and Randerson, J. T.: A human-driven decline in global burned area, Science, 356, 1356–1362, https://doi.org/10.1126/science.aal4108, 2017.
 - Arpaci, A., Eastaugh, C. S., and Vacik, H.: Selecting the best performing fire weather indices for Austrian ecoregions, Theor. Appl. Climatol., 114, 393–406, https://doi.org/10.1007/s00704-013-0839-7, 2013.
- Avitabile, V., Herold, M., Heuvelink, G. B. M., Lewis, S. L., Phillips, O. L., Asner, G. P., Armston, J., Ashton, P. S., Banin, L., Bayol, N.,
 Berry, N. J., Boeckx, P., de Jong, B. H. J., DeVries, B., Girardin, C. A. J., Kearsley, E., Lindsell, J. A., Lopez-Gonzalez, G., Lucas, R.,
 Malhi, Y., Morel, A., Mitchard, E. T. A., Nagy, L., Qie, L., Quinones, M. J., Ryan, C. M., Ferry, S. J. W., Sunderland, T., Laurin, G. V.,
 Gatti, R. C., Valentini, R., Verbeeck, H., Wijaya, A., and Willcock, S.: An integrated pan-tropical biomass map using multiple reference datasets, Global Change Biol., 22, 1406–1420, https://doi.org/10.1111/gcb.13139, 2016.
- Beuchle, R., Grecchi, R. C., Shimabukuro, Y. E., Seliger, R., Eva, H. D., Sano, E., and Achard, F.: Land cover changes in the Brazilian

 Cerrado and Caatinga biomes from 1990 to 2010 based on a systematic remote sensing sampling approach, Appl. Geogr., 58, 116–127,
 https://doi.org/10.1016/j.apgeog.2015.01.017, 2015.
 - Bondeau, A., Smith, P. C., Zaehle, S., Schaphoff, S., Lucht, W., Cramer, W., Gerten, D., Lotze-Campen, H., Müller, C., Reichstein, M., and Smith, B.: Modelling the role of agriculture for the 20th century global terrestrial carbon balance, Global Change Biol., 13, 679–706, https://doi.org/10.1111/j.1365-2486.2006.01305.x, 2007.
- Broyden, C. G.: The Convergence of a Class of Double-rank Minimization Algorithms 1. General Considerations, IMA J. Appl. Math., 6, 76–90, https://doi.org/10.1093/imamat/6.1.76, 1970.
 - Chambers, J. Q. and Artaxo, P.: Biosphere–atmosphere interactions: Deforestation size influences rainfall, Nat. Clim. Change, 7, 175, https://doi.org/10.1038/nclimate3238, 2017.
- Christian, H. J., Blakeslee, R. J., Boccippio, D. J., Boeck, W. L., Buechler, D. E., Driscoll, K. T., Goodman, S. J., Hall, J. M., Koshak, W. J.,
 Mach, D. M., and Stewart, M. F.: Global frequency and distribution of lightning as observed from space by the Optical Transient Detector,
 J. Geophys. Res. Atmos., 108, ACL4–1–ACL4–15, https://doi.org/10.1029/2002JD002347, 2003.
 - Chuvieco, E., Aguado, I., Yebra, M., Nieto, H., Salas, J., Martín, M. P., Vilar, L., Martínez, J., Martín, S., Ibarra, P., de la Riva, J., Baeza, J., Rodríguez, F., Molina, J. R., Herrera, M. A., and Zamora, R.: Development of a framework for fire risk assessment using remote sensing and geographic information system technologies, Ecol. Modell., 221, 46–58, https://doi.org/10.1016/j.ecolmodel.2008.11.017, 2010.
- 30 Civil, O. and Environmental Engineering/Princeton University, D.: Global Meteorological Forcing Dataset for Land Surface Modeling, UCAR/NCAR - Research Data Archive, https://rda.ucar.edu/datasets/ds314.0, 2006.
 - Cochrane, M. A. and Laurance, W. F.: Synergisms among Fire, Land Use, and Climate Change in the Amazon, Ambio, 37, 522–527, https://doi.org/10.2307/25547943, 2008.
- Fader, M., Rost, S., Müller, C., Bondeau, A., and Gerten, D.: Virtual water content of temperate cereals and maize: Present and potential future patterns, J. Hydrol., 384, 218–231, https://doi.org/10.1016/j.jhydrol.2009.12.011, 2010.
 - Fletcher, R.: A new approach to variable metric algorithms, comjnl., 13, 317–322, https://doi.org/10.1093/comjnl/13.3.317, 1970.

- Forkel, M., Carvalhais, N., Schaphoff, S., V. Bloh, W., Migliavacca, M., Thurner, M., and Thonicke, K.: Identifying environmental controls on vegetation greenness phenology through model—data integration, Biogeosciences, 11, 7025–7050, https://doi.org/10.5194/bg-11-7025-2014, 2014.
- Forkel, M., Dorigo, W., Lasslop, G., Teubner, I., Chuvieco, E., and Thonicke, K.: A data-driven approach to identify controls on global fire activity from satellite and climate observations (SOFIA V1), Geosci. Model Dev., 10, 4443–4476, https://doi.org/10.5194/gmd-10-4443-2017, 2017.
 - Forkel, M., Andela, N., Harrison, S. P., Lasslop, G., Marle, M. v., Chuvieco, E., Dorigo, W., Forrest, M., Hantson, S., Heil, A., Li, F., Melton, J., Sitch, S., Yue, C., and Arneth, A.: Emergent relationships with respect to burned area in global satellite observations and fire-enabled vegetation models, Biogeosciences, 16, 57–76. https://doi.org/10.5194/bg-16-57-2019, 2019a.
- Forkel, M., Dorigo, W., Lasslop, G., Chuvieco, E., Hantson, S., Heil, A., Teubner, I., Thonicke, K., and Harrison, S. P.: Recent global and regional trends in burned area and their compensating environmental controls, Environ. Res. Commun., 1, 051005, https://doi.org/10.1088/2515-7620/ab25d2, 2019b.
 - Gerten, D., Schaphoff, S., Haberlandt, U., Lucht, W., and Sitch, S.: Terrestrial vegetation and water balance—hydrological evaluation of a dynamic global vegetation model, J. Hydrol., 286, 249–270, https://doi.org/10.1016/j.jhydrol.2003.09.029, 2004.
- 15 Giglio, L., Randerson, J. T., and van der Werf, G. R.: Analysis of daily, monthly, and annual burned area using the fourth-generation global fire emissions database (GFED4), J. Geophys. Res. Biogeosci., 118, 317–328, https://doi.org/10.1002/jgrg.20042, 2013.
 - Goff, J. and Gratch, S.: List 1947, Smithsonian meteorological tables, Transactions of the American Society of Ventilation Engineering, 52, 95, 1946.
 - Goldewijk, K. K., Beusen, A., van Drecht, G., and de Vos, M.: The HYDE 3.1 spatially explicit database of human-induced global land-use change over the past 12,000 years, Global Ecol. Biogeogr., 20, 73–86, https://doi.org/10.1111/j.1466-8238.2010.00587.x, 2011.
 - Goldfarb, D.: A family of variable-metric methods derived by variational means, Math. Comput., 24, 23–26, https://doi.org/10.1090/S0025-5718-1970-0258249-6, 1970.
 - Gupta, H. V., Kling, H., Yilmaz, K. K., and Martinez, G. F.: Decomposition of the mean squared error and NSE performance criteria: Implications for improving hydrological modelling, J. Hydrol., 377, 80–91, https://doi.org/10.1016/j.jhydrol.2009.08.003, 2009.
- 25 Hantson, S., Arneth, A., Harrison, S. P., Kelley, D. I., Prentice, I. C., Rabin, S. S., Archibald, S., Mouillot, F., Arnold, S. R., Artaxo, P., Bachelet, D., Ciais, P., Forrest, M., Friedlingstein, P., Hickler, T., Kaplan, J. O., Kloster, S., Knorr, W., Laslop, G., Li, F., Melton, J. R., Meyn, A., Sitch, S., Spessa, A., van der Werf, G. R., Voulgarakis, A., and Yue, C.: The status and challenge of global fire modelling, Copernicus Publications, http://centaur.reading.ac.uk/53498, 2016.
 - Harvard: Harvard WorldMap, http://worldmap.harvard.edu, [Online; accessed 27. Mar. 2019], 2019.

- Hoffmann, W. A., Jackson, R. B., Hoffmann, W. A., and Jackson, R. B.: Vegetation–Climate Feedbacks in the Conversion of Tropical Savanna to Grassland, https://doi.org/10.1175/1520-0442(2000)013<1593:VCFITC>2.0.CO;2, [Online; accessed 28. Mar. 2019], 2000.
 - IBGE: Mapa de Biomas e de Vegetação, https://ww2.ibge.gov.br/home/presidencia/noticias/21052004biomashtml.shtm, [Online; accessed 14. Feb. 2019], 2019.
- Jolly, W. M., Cochrane, M. A., Freeborn, P. H., Holden, Z. A., Brown, T. J., Williamson, G. J., and Bowman, D. M. J. S.: Climate-induced variations in global wildfire danger from 1979 to 2013, Nat. Commun., 6, 7537, https://doi.org/10.1038/ncomms8537, 2015.
 - Keeley, J. E., Pausas, J. G., Rundel, P. W., Bond, W. J., and Bradstock, R. A.: Fire as an evolutionary pressure shaping plant traits, Trends Plant Sci., 16, 406–411, https://doi.org/10.1016/j.tplants.2011.04.002, 2011.

- Keenan, T. F., Carbone, M. S., Reichstein, M., and Richardson, A. D.: The model-data fusion pitfall: assuming certainty in an uncertain world, Oecologia, 167, 587, https://doi.org/10.1007/s00442-011-2106-x, 2011.
- Keetch, J. J. and Byram, G. M.: A Drought Index for Forest Fire Control, Res. Pap. SE-38. Asheville, NC: U.S. Department of Agriculture, Forest Service, Southeastern Forest Experiment Station. 35 p. 038, https://www.fs.usda.gov/treesearch/pubs/40, 1968.
- Kelley, D. I., Prentice, I. C., Harrison, S. P., Wang, H., Simard, M., Fisher, J. B., and Willis, K. O.: A comprehensive benchmarking system for evaluating global vegetation models, Biogeosciences, 10, 3313, https://doi.org/10.5194/bg-10-3313-2013, 2013.
 - Knorr, W., Arneth, A., and Jiang, L.: Demographic controls of future global fire risk, Nat. Clim. Change, 6, 781, https://doi.org/10.1038/nclimate2999, 2016.
- Krawchuk, M. A. and Moritz, M. A.: Constraints on global fire activity vary across a resource gradient, Ecology, 92, 121–132, https://doi.org/10.1890/09-1843.1, 2011.
 - Lahsen, M., Bustamante, M. M. C., and Dalla-Nora, E. L.: Undervaluing and Overexploiting the Brazilian Cerrado at Our Peril, Environment: Science and Policy for Sustainable Development, 58, 4–15, https://doi.org/10.1080/00139157.2016.1229537, 2016.
 - Langmann, B., Duncan, B., Textor, C., Trentmann, J., and van der Werf, G. R.: Vegetation fire emissions and their impact on air pollution and climate, Atmos. Environ., 43, 107–116, https://doi.org/10.1016/j.atmosenv.2008.09.047, 2009.
- 15 Lasslop, G., Thonicke, K., and Kloster, S.: SPITFIRE within the MPI Earth system model: Model development and evaluation, J. Adv. Model. Earth Syst., 6, 740–755, https://doi.org/10.1002/2013MS000284, 2014.
 - Lasslop, G., Hantson, S., and Kloster, S.: Influence of wind speed on the global variability of burned fraction: a global fire model's perspective, Int. J. Wildland Fire, 24, 989–1000, https://doi.org/10.1071/WF15052, 2015.
 - Lasslop, G., Brovkin, V., Reick, C. H., Bathiany, S., and Kloster, S.: Multiple stable states of tree cover in a global land surface model due to a fire-vegetation feedback, Geophys. Res. Lett., 43, 6324–6331, https://doi.org/10.1002/2016GL069365, 2016.

25

- Li, W., MacBean, N., Ciais, P., Defourny, P., Lamarche, C., Bontemps, S., Houghton, R. A., and Peng, S.: Gross and net land cover changes in the main plant functional types derived from the annual ESA CCI land cover maps (1992–2015), Earth Syst. Sci. Data, 10, 219–234, https://doi.org/10.5194/essd-10-219-2018, 2018.
- Mebane, Jr., W. R. and Sekhon, J. S.: Genetic Optimization Using Derivatives: The rgenoud Package for R, Journal of Statistical Software, 42, 1–26, https://doi.org/10.18637/jss.v042.i11, 2011.
- Moreira de Araújo, F., Ferreira, L. G., and Arantes, A. E.: Distribution Patterns of Burned Areas in the Brazilian Biomes: An Analysis Based on Satellite Data for the 2002–2010 Period, Remote Sens., 4, 1929–1946, https://doi.org/10.3390/rs4071929, 2012.
- Nachtergaele, F. O., van Velthuizen, H. T., and Verelst, L.: Harmonized World Soil Database, http://pure.iiasa.ac.at/id/eprint/8958, [Online; accessed 4. Oct. 2018], 2009.
- Panisset, J. S., Libonati, R., Gouveia, C. M. P., Machado-Silva, F., França, D. A., França, J. R. A., and Peres, L. F.: Contrasting patterns of the extreme drought episodes of 2005, 2010 and 2015 in the Amazon Basin, Int. J. Climatol., 38, 1096–1104, https://doi.org/10.1002/joc.5224, 2017.
 - Parente, L., Ferreira, L., Faria, A., Nogueira, S., Araújo, F., Teixeira, L., and Hagen, S.: Monitoring the brazilian pasturelands: A new mapping approach based on the landsat 8 spectral and temporal domains, Int. J. Appl. Earth Obs. Geoinf., 62, 135–143, https://doi.org/10.1016/j.jag.2017.06.003, 2017.
 - Pechony, O. and Shindell, D. T.: Fire parameterization on a global scale, J. Geophys. Res. Atmos., 114, https://doi.org/10.1029/2009JD011927, 2009.

- Pfeiffer, M., Spessa, A., and Kaplan, J. O.: A model for global biomass burning in preindustrial time: LPJ-LMfire (v1.0), Geosci. Model Dev., 6, 643–685, https://doi.org/10.5194/gmd-6-643-2013, 2013.
- Prado, D. E.: As caatingas da América do Sul, Ecologia e conservação da Caatinga, 2, 3-74, 2003.
- Pyne, S. J., Andrews, P. L., and Laven, R. D.: Introduction to wildland fire, 2nd edition revised, DigitalCommons@USU, https: //digitalcommons.usu.edu/barkbeetles/135, 1996.
 - Quéré, C. L., Moriarty, R., Andrew, R. M., Canadell, J. G., Sitch, S., Korsbakken, J. I., Friedlingstein, P., Peters, G. P., Andres, R. J., Boden, T. A., Houghton, R. A., House, J. I., Keeling, R. F., Tans, P., Arneth, A., Bakker, D. C. E., Barbero, L., Bopp, L., Chang, J., Chevallier, F., Chini, L. P., Ciais, P., Fader, M., Feely, R. A., Gkritzalis, T., Harris, I., Hauck, J., Ilyina, T., Jain, A. K., Kato, E., Kitidis, V., Goldewijk, K. K., Koven, C., Landschützer, P., Lauvset, S. K., Lefèvre, N., Lenton, A., Lima, I. D., Metzl, N., Millero, F., Munro, D. R., Murata,
- A., Nabel, J. E. M. S., Nakaoka, S., Nojiri, Y., O'Brien, K., Olsen, A., Ono, T., Pérez, F. F., Pfeil, B., Pierrot, D., Poulter, B., Rehder, G., Rödenbeck, C., Saito, S., Schuster, U., Schwinger, J., Séférian, R., Steinhoff, T., Stocker, B. D., Sutton, A. J., Takahashi, T., Tilbrook, B., van der Laan-Luijkx, I. T., van der Werf, G. R., van Heuven, S., Vandemark, D., Viovy, N., Wiltshire, A., Zaehle, S., and Zeng, N.: Global Carbon Budget 2015, Earth Syst. Sci. Data, 7, 349–396, https://doi.org/10.5194/essd-7-349-2015, 2015.
- Rabin, S. S., Ward, D. S., Malyshev, S. L., Magi, B. I., Shevliakova, E., and Pacala, S. W.: A fire model with distinct crop, pasture, and non-agricultural burning: use of new data and a model-fitting algorithm for FINAL.1, Geosci. Model Dev., 11, 815–842, https://doi.org/10.5194/gmd-11-815-2018, 2018.
 - Ray, D., Nepstad, D., and Moutinho, P.: MICROMETEOROLOGICAL AND CANOPY CONTROLS OF FIRE SUSCEPTIBILITY IN A FORESTED AMAZON LANDSCAPE, Ecol. Appl., 15, 1664–1678, https://doi.org/10.1890/05-0404, 2005.
- Rodell, M., Houser, P. R., Jambor, U., Gottschalck, J., Mitchell, K., Meng, C.-J., Arsenault, K., Cosgrove, B., Radakovich, J., Bosilovich,
 M., Entin, J. K., Walker, J. P., Lohmann, D., Toll, D., Rodell, M., Houser, P. R., Jambor, U., Gottschalck, J., Mitchell, K., Meng, C.-J.,
 Arsenault, K., Cosgrove, B., Radakovich, J., Bosilovich, M., Entin, J. K., Walker, J. P., Lohmann, D., and Toll, D.: The Global Land Data
 Assimilation System, Bull. Am. Meteorol. Soc., https://doi.org/10.1175/BAMS-85-3-381, 2004.
 - Rogers, B. M., Soja, A. J., Goulden, M. L., and Randerson, J. T.: Influence of tree species on continental differences in boreal fires and climate feedbacks, Nat. Geosci., 8, 228, https://doi.org/10.1038/ngeo2352, 2015.
- Roitman, I., Bustamante, M. M. C., Haidar, R. F., Shimbo, J. Z., Abdala, G. C., Eiten, G., Fagg, C. W., Felfili, M. C., Felfili, J. M., Jacobson, T. K. B., Lindoso, G. S., Keller, M., Lenza, E., Miranda, S. C., Pinto, J. R. R., Rodrigues, A. A., Delitti, W. B. C., Roitman, P., and Sampaio, J. M.: Optimizing biomass estimates of savanna woodland at different spatial scales in the Brazilian Cerrado: Re-evaluating allometric equations and environmental influences, PLoS One, 13, e0196742, https://doi.org/10.1371/journal.pone.0196742, 2018.
 - Rothermel, R. C.: A mathematical model for predicting fire spread in wildland fuels, DigitalCommons@USU, Res. Pap. INT-115, 1, https://digitalcommons.usu.edu/barkbeetles/438, 1972.

- Schaphoff, S., Heyder, U., Ostberg, S., Gerten, D., Heinke, J., and Lucht, W.: Contribution of permafrost soils to the global carbon budget, Environ. Res. Lett., 8, 014 026, https://doi.org/10.1088/1748-9326/8/1/014026, 2013.
- Schaphoff, S., Bloh, W. v., Rammig, A., Thonicke, K., Biemans, H., Forkel, M., Gerten, D., Heinke, J., Jägermeyr, J., Knauer, J., Langerwisch, F., Lucht, W., Müller, C., Rolinski, S., and Waha, K.: LPJmL4 a dynamic global vegetation model with managed land Part 1: Model description, Geosci. Model Dev., 11, 1343–1375, https://doi.org/10.5194/gmd-11-1343-2018, 2018a.
- Schaphoff, S., Forkel, M., Müller, C., Knauer, J., Bloh, W. v., Gerten, D., Jägermeyr, J., Lucht, W., Rammig, A., Thonicke, K., and Waha, K.: LPJmL4 a dynamic global vegetation model with managed land Part 2: Model evaluation, Geosci. Model Dev., 11, 1377–1403, https://doi.org/10.5194/gmd-11-1377-2018, 2018b.

- Seager, R., Hooks, A., Williams, A. P., Cook, B., Nakamura, J., Henderson, N., Seager, R., Hooks, A., Williams, A. P., Cook, B., Nakamura, J., and Henderson, N.: Climatology, Variability, and Trends in the U.S. Vapor Pressure Deficit, an Important Fire-Related Meteorological Quantity*, https://journals.ametsoc.org/doi/abs/10.1175/JAMC-D-14-0321.1, [Online; accessed 19. Oct. 2018], 2015.
- Sedano, F. and Randerson, J.: Multi-scale influence of vapor pressure deficit on fire ignition and spread in boreal forest ecosystems, https://doi.org/10.5194/bg-11-3739-2014, [Online; accessed 11. Oct. 2018], 2014.

10

15

20

- Shanno, D. F.: Conditioning of quasi-Newton methods for function minimization, Math. Comput., 24, 647–656, https://doi.org/10.1090/S0025-5718-1970-0274029-X, 1970.
- Silvério, D. V., Brando, P. M., Balch, J. K., Putz, F. E., Nepstad, D. C., Oliveira-Santos, C., and Bustamante, M. M. C.: Testing the Amazon savannization hypothesis: fire effects on invasion of a neotropical forest by native cerrado and exotic pasture grasses, Philosophical Transactions of the Royal Society B: Biological Sciences, https://royalsocietypublishing.org/doi/full/10.1098/rstb.2012.0427, 2013.
- Sitch, S., Smith, B., Prentice, I. C., Arneth, A., Bondeau, A., Cramer, W., Kaplan, J. O., Levis, S., Lucht, W., Sykes, M. T., Thonicke, K., and Venevsky, S.: Evaluation of ecosystem dynamics, plant geography and terrestrial carbon cycling in the LPJ dynamic global vegetation model, Global Change Biol., 9, 161–185, https://doi.org/10.1046/j.1365-2486.2003.00569.x, 2003.
- Thonicke, K., Venevsky, S., Sitch, S., and Cramer, W.: The role of fire disturbance for global vegetation dynamics: coupling fire into a Dynamic Global Vegetation Model, Global Ecol. Biogeogr., 10, 661–677, https://doi.org/10.1046/j.1466-822X.2001.00175.x, 2001.
- Thonicke, K., Spessa, A., Prentice, I. C., Harrison, S. P., Dong, L., and Carmona-Moreno, C.: The influence of vegetation, fire spread and fire behaviour on biomass burning and trace gas emissions: results from a process-based model, Biogeosciences, 7, 1991–2011, https://doi.org/10.5194/bg-7-1991-2010, 2010.
- van der Werf, G. R., Dempewolf, J., Trigg, S. N., Randerson, J. T., Kasibhatla, P. S., Giglio, L., Murdiyarso, D., Peters, W., Morton, D. C., Collatz, G. J., Dolman, A. J., and DeFries, R. S.: Climate regulation of fire emissions and deforestation in equatorial Asia, Proc. Natl. Acad. Sci. U.S.A., 105, 20350–20355, https://doi.org/10.1073/pnas.0803375105, 2008.
- Venevsky, S., Thonicke, K., Sitch, S., and Cramer, W.: Simulating fire regimes in human-dominated ecosystems: Iberian Peninsula case study, Global Change Biol., 8, 984–998, https://doi.org/10.1046/j.1365-2486.2002.00528.x, 2002.
- Wagner, C. E. V., Forest, P., Station, E., Ontario, C. R., Francais, R. U. E., and Davis, H. J.: Development and Structure of the Canadian Forest FireWeather Index System, Can. For. Serv., Forestry Tech. Rep, 1987.
- Werf, G. R. v. d., Randerson, J. T., Giglio, L., Collatz, G. J., Mu, M., Kasibhatla, P. S., Morton, D. C., DeFries, R. S., Jin, Y., and Leeuwen, T. v.: Global fire emissions and the contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009), Atmos. Chem. Phys., 10, 11707–11735, https://doi.org/10.5194/acp-10-11707-2010, 2010.
- Werf, G. R. v. d., Randerson, J. T., Giglio, L., Leeuwen, T. T. v., Chen, Y., Rogers, B. M., Mu, M., Marle, M. J. E. v., Morton, D. C., Collatz, G. J., Yokelson, R. J., and Kasibhatla, P. S.: Global fire emissions estimates during 1997–2016, Earth Syst. Sci. Data, 9, 697–720, https://doi.org/10.5194/essd-9-697-2017, 2017.
 - Willmott, C. J.: Some Comments on the Evaluation of Model Performance, Bull. Am. Meteorol. Soc., https://doi.org/10.1175/1520-0477(1982)063<1309:SCOTEO>2.0.CO;2, 1982.
- Wilson, R. A.: A reexamination of fire spread in free-burning porous fuel beds /, http://www.sidalc.net/cgi-bin/wxis.exe/?IsisScript= COLPOS.xis&method\=post&formato=2&cantidad=1&expresion=mfn=001388, [Online; accessed 4. Oct. 2018], 1982.
 - Yue, X. and Unger, N.: Fire air pollution reduces global terrestrial productivity, Nat. Commun., 9, 5413, https://doi.org/10.1038/s41467-018-07921-4, 2018.