Fast domain-aware neural network emulation of a planetary boundary layer parameterization in a numerical weather forecast model

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Abstract. Parameterizations for physical processes in weather and climate models are 11 computationally expensive. We use model output from a set of simulations performed using the 12 Weather Research Forecast (WRF) model to train deep neural networks and evaluate whether 13 14 trained models can provide an accurate alternative to the physics-based parameterizations. Specifically, we develop an emulator using deep neural networks for a planetary boundary layer 15 (PBL) parameterization in the WRF model. PBL parameterizations are used in atmospheric 16 models to represent the diurnal variation of the formation and collapse of the atmospheric 17 boundary layer-the lowest part of the atmosphere. The dynamics and turbulence, as well as 18 velocity, temperature, and humidity profiles within the boundary layer are all critical for 19 20 determining many of the physical processes in the atmosphere. PBL parameterizations are used 21 to represent these processes that are usually unresolved in a typical numerical weather model that operates at horizontal spatial scales in the tens of kilometers. We demonstrate that a domain-22 23 aware deep neural network, which takes account of underlying domain structure (e.g., nonlocal 24 mixing between multiple vertical layers), can successfully simulate the vertical profiles within the boundary layer of velocities, temperature, and water vapor over the entire diurnal cycle. We 25 26 then assess the spatial transferability of the domain-aware neural networks by using a trained 27 model from one location to nearby locations. Results show that a single trained model from a location over the midwestern United States produces predictions of wind speed, temperature, and 28 water vapor profiles over the entire diurnal cycle and nearby locations with similar terrain 29 30 conditions with correlations higher than 0.9 when compared with the WRF simulations used as 31 the training dataset.

32 **1 Introduction**

Model developers use approximations to represent the physical processes involved in climate and 33 weather that cannot be resolved at the spatial resolution of the model grids or in cases where the 34 phenomena are not fully understood (Williams, 2005). These approximations are referred to as 35 parameterizations (McFarlane, 2011). While these parameterizations are designed to be 36 computationally efficient, calculation of a model physics package still takes a good portion of the 37 total computational time. For example, in the community atmospheric model (CAM) developed 38 by National Center for Atmospheric Research (NCAR), with spatial resolution of approximately 39 300 km and 26 vertical levels, the physical parameterizations account for about 70% of the total 40 computational burden (Krasnopolsky and Fox-Rabinovitz, 2006). In the Weather Research 41 Forecast (WRF) model, with spatial resolution of tens of kilometers, time spent by physics is 42 43 approximately 40% of the computational burden. The input and output overhead is around 20% of the computational time at low node count (100's) and can increase significantly at higher node 44 count as a percentage of the total wall-clock time. 45

An increasing need in the climate community is performing high spatial resolution simulations (grid spacing of 4 km or less) and generating large ensembles of these simulations in order to address uncertainty in the model projections and to assess risk and vulnerability due to climate variability at local scale. Developing process emulators (Leeds et al., 2013; Lee et al., 2011) that can reduce the time spent in calculating the physical processes will lead to much faster model simulations, enabling researchers to generate high spatial resolution simulations and a large number of ensemble members.

A neural network (NN) is composed of multiple layers of simple computational modules, where 53 54 each module transforms its inputs to a nonlinear output. Given sufficient data, an appropriate NN can model the underlying nonlinear functional relationship between inputs and outputs with 55 56 minimal human effort. During the past two decades, NN techniques have found a variety of applications in atmospheric science. For example, Collins and Tossot (2015) developed an 57 58 artificial NN model by taking numerical weather prediction model (e.g., WRF) output as input to predict thunderstorm occurrence within a few hundreds of square kilometers about 12 hours in 59 60 advance. Krasnopolsky et al. (2016) used NN techniques for filling the gaps in satellite measurements of ocean color data. Scher (2018) used deep learning to emulate the complete 61

physics and dynamics of a simple general circulation model and indicated a potential capability 62 63 of weather forecasts using this NN-based emulator. Neural networks are particularly appealing for emulations of model physics parameterizations in numerical weather and climate modeling, 64 where the goal is to find nonlinear functional relationship between inputs and outputs (Cybenko, 65 1989; Hornik, 1991; Chen and Chen, 1995a,b; Attali and Pagès, 1997). NN techniques can be 66 applied to weather and climate modeling in two ways. One approach involves developing new 67 parameterizations by using NNs. For example, Chevallier et al. (1998; 2000) developed a new 68 69 NN-based longwave radiation parameterization, NeuroFlux, which has been used operationally in the European Centre for Medium-Range Weather Forecasts four-dimensional variational data 70 assimilation system. NeuroFlux is found eight times faster than the previous parameterization. 71 Krasnopolsky et al. (2013) developed a stochastic convection parameterization based on learning 72 73 from data simulated by a cloud-resolving model (CRM), initialized with and forced by the observed meteorological data. The NN convection parameterization was tested in the NCAR 74 75 CAM and produced reasonable and promising results for the tropical Pacific region. Jiang et al. (2018) developed a deep NN-based algorithm or parameterization to be used in the WRF model 76 77 to provide flow-dependent typhoon-induced sea surface temperature cooling. Results based on four typhoon case studies showed that the algorithm reduced maximum wind intensity error by 78 79 60–70% compared with using the WRF model. The other approach for applying NN to weather and climate modeling is to emulate existing parameterizations in these models. For example, 80 81 Krasnopolsky et al. (2005) developed an NN-based emulator for imitating an existing atmospheric longwave radiation parameterization for the NCAR CAM. They used output from 82 the CAM simulations with the original parameterization for the NN training. They found the 83 NN-based emulator was 50–80 times faster than the original parameterization and produced 84 85 almost identical results.

We study NN models to emulate existing physical parameterizations in atmospheric models. Process emulators that can reproduce physics parameterization can ultimately lead to the development of a faster model emulator that can operate at very high spatial resolution as compared with most current model emulators that have tended to focus on simplified physics (Kheshigi et al., 1999). Specifically, this study involves the design and development of a domain-aware NN to emulate a PBL parameterization using 22-year-long output created by a set of WRF simulations. To the best of our knowledge, we are among the first to apply deep neural

networks to the WRF model to explore the emulation of physics parameterizations. As far as we 93 know from the literature available at the time of this writing, the only application of NNs for 94 emulating the parameterizations in the WRF model is by Krasnopolsky et al. (2017). In their 95 study, a three-layer NN was trained to reproduce the behavior of the Thompson microphysics 96 (Thompson 2008) scheme in the WRF-ARW model. While we focus on learning the PBL 97 parameterization and developing domain-aware NN for emulation of PBL, the ultimate goal of 98 our on-going project is to build an NN-based algorithm to empirically understand the process in 99 the numerical weather/climate models that could be used to replace the physics parameterizations 100 that were derived from observational studies. This emulated model would be computationally 101 efficient, making the generation of large ensemble simulations feasible at very high 102 spatial/temporal resolutions with limited computational resources. The key objectives of this 103 104 study are to answer the following questions specifically for PBL parameterization emulation: (1) What and how much data do we need to train the model? (2) What type of NN should we apply 105 106 for the PBL parameterization studied here? (3) Is the NN emulator accurate compared with the original physical parameterization? This paper is organized as follows. Section 2 describes the 107 108 data and the neural network developed in this study. The efficacy of the neural network is 109 investigated in Section 3. Discussion and summary follow in Section 4.

110 2 Data and Method

111 **2.1 Data**

The data we use in this study is 22-year output from the regional climate model WRF version 112 3.3.1, driven by NCEP-R2 for the period 1984-2005. WRF is a fully compressible, 113 nonhydrostatic, regional numerical prediction system with proven suitability for a broad range of 114 applications. The WRF model configuration and evaluations are given by Wang and Kotamarthi 115 (2014). Covering all the troposphere are 38 vertical layers, between the surface to approximately 116 16 km (100 hPa). The lowest 17 layers cover from the surface to about 2 km above the ground. 117 The PBL parameterization we used for this WRF simulation is known as the YSU scheme (Hong 118 et al., 2006). The YSU scheme uses a nonlocal-mixing scheme with an explicit treatment of 119 120 entrainment at the top of the boundary layer and a first-order closure for the Reynolds-averaged turbulence equations of momentum of air within the PBL. 121

The goal of the work described here is to develop an NN-based parameterization emulator that can be used to replace the PBL parameterization in the WRF model. Thus, we expect the NN submodel to receive a set of inputs that are equivalent to the inputs provided to the YSU scheme at each timestep.

126 Table 1 shows the architecture in terms of inputs and outputs used in our experiments. The inputs 127 are near-surface characteristics including 2-meter water vapor and air temperature, 10-meter zonal and meridional wind, ground heat flux, incoming shortwave radiation, incoming longwave 128 radiation, PBL height, sensible heat flux, latent heat flux, surface friction velocity, ground 129 temperature, soil temperature at 2 m below the ground, soil moisture at 0-0.3cm below the 130 131 ground, and geostrophic wind component at 700 hPa. The outputs for the NN architecture are the vertical profiles of the following five model prognostic and diagnostic fields: temperature, water 132 133 vapor mixing ratio, zonal and meridional wind (including speed and direction), as well as vertical motions. In this study we develop an NN emulation of the PBL parameterization; hence we focus 134 135 only on predicting the profiles within the PBL, which is on average around 200 m and 400 m during the night and afternoon of winter, respectively, and around 400 m and 1300 m during the 136 137 night and afternoon of summer, respectively, for the locations studied here. The middle and upper troposphere (all layers above the PBL) are considered fully resolved by the dynamics 138 139 simulated by the model and hence not parameterized. Therefore, we do not consider the levels 140 above PBL height because (1) they carry no information about input/output functional dependence that affects the PBL and (2) if not removed, they introduce additional noise in the 141 training. Specifically, we use the WRF output from the first 17 layers, which are within 1,900 142 meters and well cover the PBL. 143

144 **2.2 Deep neural networks for PBL parameterization emulation**

A class of machine learning approaches that is particularly suitable for emulation of PBL parameterization is supervised learning. This approach models the relationship between the outputs and independent input variables by using training data (x_i , y_i), for $x_i \in T \subset D$, where T is a set of training points, D is the full data set, and x_i and $y_i = f(x_i)$ are inputs and its corresponding output y_i , respectively. The function *f* that maps the inputs to the outputs is typically unknown and hard to derive analytically. The goal of the supervised learning approach is to find a surrogate function *h* for *f* such that the difference between $f(x_i)$ and $h(x_i)$ is minimal for all $x_i \in T$. Many supervised learning algorithms exist in the machine learning literature. This study focuseson deep neural networks (DNNs).

DNNs are composed of an input layer, a series of hidden layers, and an output layer. The input 154 layer receives the input x_i, which is connected to the hidden layers. Each hidden layer receives 155 156 inputs from the previous hidden layer (except the first hidden layer that is connected to the input 157 layer) and perform certain nonlinear transformations through a system of weighted connections and a nonlinear activation function on the received input values. The last hidden layer is 158 connected to the output layer from which the predicted values are obtained. The training data is 159 given to the DNN through the input neural layer. The training procedure consists of modifying 160 161 the weights of the connections in the network to minimize a user-defined objective function that measures the prediction error of the network. Each iteration of the training procedure comprises 162 163 two phases: forward pass and backward pass. In the forward pass, the training data are passed to the network and the prediction error is computed; in the backward pass, the gradients of the error 164 165 function with respect to all the weights in the network is computed and used to update the weights in order to minimize the error. Once the entire dataset pass both forward and backward 166 167 through the DNN (with many iterations), one epoch is completed.

168 **Deep feed-forward neural network (FFN):** This is a fully connected feed-forward DNN 169 constructed as a sequence of *K* hidden layers, where the input of the *i*th hidden layer is from {*i*-170 1}th hidden layer and the output of the *i*th hidden layer is given as the input of the {*i*+1}th 171 hidden layer. The sizes of the input and output neural layers are 16 (= near-surface variables) and 172 85 (= 17 vertical levels \times 5 output variables). See Figure 1a for an illustration.

173 While the FFN is a typical way of applying NN for finding the nonlinear relationship between 174 input and output, a key drawback is that it does not consider the underlying PBL structure, such as the vertical connection between different vertical levels within the PBL. In fact, the FFN does 175 176 not know which data (among the 85 variables) belongs to which vertical levels in a certain 177 profile. This is not typically needed for NNs in general and in fact is usually avoided because, for classification and regression, one can find visual features regardless of their locations. For 178 example, a picture can be classified as a certain object even that object has never appeared in the 179 180 given location in the training set. In our case, however, the location is fixed and the profiles over that location is distinguishable from other locations if they have different terrain conditions. 181

Consequently, it is desired to learn vertical connection between multiple layers within the PBL in the forecast. For example, the feature at a lower level of a profile plays a role in the feature at a higher level and can help refine the output at the higher level and accordingly the entire profile. This dependence may inform the NN and provide better accuracy and data efficiency. To that end, we develop two variants of DNNs for PBL emulation.

187 Hierarchically connected network with previous layer only connection (HPC): We assume that the outputs at each altitude level depend not only on the 16 near-surface variables but also 188 on the *adjacent* altitude level below it. To model this explicitly, we develop a DNN variant as 189 follows: the input layer is connected to the first hidden layer followed by the output layer of size 190 191 5 (five variable at each layer: temperature, water vapor, zonal and meridional wind, and vertical motions) that corresponds to the first PBL. This output layer along with the input layer is 192 193 connected to second hidden layer, which is connected to the second output layer of size 5 that corresponds to the second PBL. Thus, the input to an *i*th hidden layer comprises the input layer 194 195 of the 16 near-surface variables and the *i*-1th output layer below it See Figure 1b for an example.

Hierarchically connected network with all previous layers connection (HAC): We assume that the outputs at each PBL depend not only on the 16 near-surface variables but also on *all* altitude levels below it. To model this explicitly, we modify HPC DNN as follows: the input to an *i*th hidden layer comprises the input layer of the 16 near-surface variables and *all* output layers $\{1, 2, ..., i-1\}$ below it. See Figure 1c for an example.

From the physical process perspective, HPC and HAC considers both local and non-local mixing processes within the PBL by taking into account not only the connection between a given point and its adjacent point (local mixing), but also the connections from multiple vertical altitude levels (e.g., surface and all the points that below the given points). Compared with solely local mixing process, non-local mixing process is showed to perform more accurately in simulating deeper mixing within an unstable PBL (Cohen et al. 2015).

From the neural network perspective, the key advantage of HPC and HAC over FNN is effective back-propagation while training. In HPC and HAC, each hidden layer has an output layer; consequently, during the back propagation, the gradients from each of the output layer can be used to update the weights of the hidden layer directly to minimize the error for PBL specific outputs.

212 **2.3 Setup**

For preprocessing, we applied StandardScaler (removes the mean and scales each variable to unit variance) and MinMaxScaler (scales each variable between 0 and 1) transformations before training, and we applied the inverse transformation after prediction so that the evaluation metrics are computed on the original scale.

We note that there is no default value for *N* units in a dense hidden layer. We conducted an experimental study on FFN and found that setting *N* to 16 results in good predictions. Therefore, we used the same value of N = 16 in HPC and HAC.

For the implementation of DNN, we used Keras (version 2.0.8), a high-level neural network Python library that runs on the top of the TensorFlow library (version 1.3.0). We used the scikitlearn library (version 0.19.0) for the preprocessing module. The experiments were run on a Python (Intel distribution, version 3.6.3) environment.

All three DNNs used the following setup for training: optimizer = adam, learning rate = 0.001, epochs = 1000, batch size = 64. Note that batch size and number of epochs define the number of randomly sampled training points required before updating the model parameters and the number times that training will work through the entire training dataset. To avoid overfitting issues in DNNs, we use an early stopping criterion in which the training stops when the validation error does not reduce for 10 subsequent epochs.

The 22-year data from the WRF simulation was partitioned into three parts: a training set consisting of 20 years (1984–2003) of 3-hourly data to train the NN; a development set (also called validation set) consisting of 1 year (2004) of 3-hourly data used to tune the algorithm's hyperparameters and to control overfitting (the situation where the trained network predicts well on the training data but not on the test data); and a test set consisting of 1 year of records (2005) for prediction and evaluations.

We ran training and inference on a NVIDIA DGX-1 platform: Dual 20-Core Intel Xeon E5-2698
v4 2.2 GHz processor with 8 NVIDIA P100 GPUs with 512 GB of memory. The DNN's training
and inference leveraged only a single GPU.

239 **3 Results**

In the following discussion we evaluate the efficacy of the three DNNs by comparing their 240 prediction results with WRF model simulations. We refer to the results of WRF model 241 242 simulations as observations because the DNN learns all the knowledge from the WRF model output, not from in situ measurements. We refer to the values from the DNN models as 243 predictions. We initiate our DNN development at one grid cell from WRF output that is close to 244 a site in the midwestern United States (Logan, Kansas, latitude= 38.8701°N; longitude= 245 100.9627°W) and another grid cell at a site in Alaska (Kenai Peninsula Borough, AK, latitude= 246 60.7237 °N; longitude=150.4484 °W) to evaluate the robustness of the developed DNNs. We 247 then apply our DNNs to an area with size of ~1100 km x 1100 km, centered at the Logan site to 248 assess the spatial transferability of the DNNs. In other words, we train our DNNs using a single 249 location, and then apply the DNNs to multiple grid points nearby. While the Alaska site has 250 251 different vertical profiles, especially for wind directions, and lower PBL heights in both January and July, the conclusion in terms of the model performance is similar to the site over Logan, 252 Kansas. 253

3.1 DNN performance in temperature and water vapor

Figure 2 shows the diurnal variation (explicitly 3 PM and 12 AM local time at Logan, Kansas) of 255 temperature and water vapor mixing ratio vertical profiles in the first 17 layers from the 256 257 observation and three DNN model predictions. The figures present results for both January and July of 2005. The dashed lines show the lowest and highest (5th and 95th percentile, respectively) 258 PBL heights for that particular time. In general, the DNNs are able to produce similar shapes of 259 260 the observed profiles, especially within the PBL. Both the temperature and water vapor mixing ratio are lower in January and higher in July. Within the PBL, the temperature and water vapor 261 do not change much with height; above the PBL to the entrainment zone, the temperature and 262 263 water vapor start decreasing. Among the three DNNs, HAC and HPC show very low bias and high accuracy in the PBL; the FFN shows a relatively large discrepancy from the observation. 264 Figure 3 shows the root-mean-square error (RMSE) and Pearson correlation coefficient (COR) 265 266 between observation and three DNN predictions in the afternoon and midnight of January and July. The RMSE and COR consider not only the time series of observation and prediction but 267 268 also their vertical profiles below the PBL heights for each particular time. Among the three DNNs, HPC and HAC always show better skill with smaller RMSEs and higher CORs than does 269

FFN. The reason is that the FFN uses only the 16 near-surface variables as inputs and all the 85 270 variables (17 layers \times 5 variables/layer) as output, and does not have the knowledge about the 271 272 vertical connections between each of the vertical levels. In contrast, HPC and HAC use both the 273 near-surface variables and the five output variables of one previous vertical level (HPC) or all previous vertical levels (HAC) as inputs for predicting a certain vertical level of each field. This 274 275 architecture is helpful for reducing errors of each hidden layer during the backward propagation. It is also important because PBL parameterizations are used to represent the vertical mixing of 276 heat, moist, and momentum within the PBL and this mixing can be across a larger scale than just 277 the adjacent altitude levels. This process is usually unresolved in a typical climate and weather 278 models that operate at horizontal spatial scales in the tens of kilometers. We find in general HAC 279 and HPC perform similarly, although in winter especially midnight when the PBL is shallow, the 280 281 RMSE of temperature predicted by HAC is larger than that predicted by HPC. In contrast, in summer especially in afternoon when the PBL is deep, the RMSE of temperature predicted by 282 HAC is smaller than that predicted by HPC. This emphasizes the importance of consideration of 283 multi-level vertical connection for deep PBL case in the DNNs. 284

3.2 DNN performance in wind component

Figure 4 shows the diurnal variation of zonal and meridional wind (including wind speed and 286 287 direction) profiles in January and July 2005 from observation and three DNN predictions. Compared with the temperature and water vapor profiles, the wind profiles are more difficult to 288 predict, especially for days (e.g., summer) that have a higher PBL. The wind direction does not 289 change much below the majority of the PBL, and it turns to westerly winds when going up and 290 beyond the PBL. The DNN prediction has difficulty predicting the profile above the PBL height, 291 as is expected because these layers are considered fully resolved by the dynamics simulated by 292 293 the WRF model and hence not parameterized. Therefore, we do not consider DNN performance at the levels above PBL height, because the DNNs carry no information about input/output 294 functional dependence that affects the PBL. The wind speed increases with height in both 295 296 January and July within the PBL. Above the PBL heights, the wind speed still increases in January but decreases in July. The reason is that in January the zonal wind, especially westerly 297 298 wind, is dominant in the atmosphere and the wind speed increases with height; in July, however, the zonal wind is relatively weak, and the meridional wind is dominant with southerly wind 299

below ~ 2 km and northerly wind above 2 km. The decrease in wind speed above the PBL is just 300 about the transition of wind direction from southerly to northerly wind. Figure 5 shows the 301 302 RMSEs and CORs between the observed and predicted wind component within the PBL. The wind component is fairly well predicted by the HAC and HPC networks with correlation above 303 0.8 for wind speed and 0.7 for wind direction except in July at midnight, which is near 0.5. 304 305 Similar to the predictions for temperature and water vapor, the FFN shows the poorest prediction accuracy with large RMSEs and low CORs, especially for wind direction in July midnight, the 306 COR is below zero. For accurately predicting the wind direction, we found that using the 307 geostrophic wind at 700 hPa as one of the inputs for the DNNs is important. 308

309 3.3 DNN dependence on length of training period

Next, we evaluate how sensitive the DNN is to the amount of available training data and how 310 much data one would need in order to train a DNN. While we present Figures 2–5 using 20-year 311 (1984–2003) training data, here we gradually decrease the length of the training set to 12 (1992– 312 313 2003), 6 (1998–2003), 2 (2002–2003) years, and 1 (2003) year. The validation data (for tuning hyper-parameters and controlling overfit) and the test data (for prediction) are kept the same as 314 in our standard training dataset, which is year 2004 and 2005, respectively. Figures 6 and 7 show 315 the RMSE and CORs between observed and predicted profiles of temperature, water vapor, and 316 317 wind component for January midnight. Overall, the FFN network depends heavily on the length of training dataset. For example, the RMSE of FFN predicted temperature decreases from 7.2 K 318 using one year of training data to 3.0 K using 20-year training data. HAC and HPC also depend 319 320 on the length of training data especially when less than 6-year training data is available, but even their worst prediction accuracy (using one year of training data) is still better than FFN using 20-321 year training data. The RMSEs of HPC and HAC predicted temperature decrease from ~2.4 K 322 using 1 year of training data to ~1.5 K using 20 years of training data. The CORs of FFN 323 predicted temperature increase from 0.73 using one year of training data to 0.92 using 20 years 324 of training data. The CORs for HPC and HAC increase slightly with more training data, but 325 326 overall they are above 0.85 using one year to 20 years of training data.

Regarding the question about how much data one would need to train a DDN, for FFN, at least from this study, the performance is not stable until one has 12 or more years of training data, which is significantly better than having only 6 years or less of training data. For HAC and HPC, however, having 6 years of training data seems sufficient to show a stable prediction. Increasing the amount of training data shows only marginal improvement in predictive accuracy. In fact, in contrast to HAC and HPC, the performance of FFN has not reached a plateau even with the 20 years of training data. This suggests that with longer training sets the predicting skill of FFN could be further improved even though it does not explicitly consider the physical process within a PBL.

336 3.4 DNN performance for nearby locations

337 This section assesses the spatial transferability of the domain-aware neural networks (specifically HAC and HPC) by using a trained model from one location (at Logan, Kansas, as presented 338 339 above) to other locations within an area with size of 1100 km x 1100 km, covering latitude from 33.431 to 44.086°N, and longitude from 107.418 to 93.6975°W, centered at the Logan site with 340 different terrain and vegetation conditions (Figure 8, top). To reduce the computational burden, 341 we pick every other 7 grid points in this area and use the 13×13 grid points (which can still 342 343 capture the terrain variability) to test the spatial transferability of the DNNs developed based on 344 the single location at Logan, Kansas. For each of the 13×13 grid points, we calculate the differences and correlations between observations and predictions. Different from the preceding 345 section, here we calculate normalized RMSEs relative to each grid point's observations averaged 346 347 over a particular time period, in order to make the comparison feasible between different grid points over the area. As shown in Figures 8 and 9 by the normalized RMSEs and Pearson 348 correlations, in general, for temperature, water vapor, and wind speed, the neural network still 349 350 work fairly well for surrounding locations and even far locations with similar terrain height, except over the grid points where the terrain height is much higher than the Logan site, and the 351 352 prediction skill gets worse with larger RMSEs. This suggests the DNNs developed based on 353 Logan site are not applicable for these locations. However, for wind direction, the prediction skill is good over the western part of the tested area, but is not so good over the far eastern part 354 of the area. One of the reasons is perhaps because that the driver of the wind direction over the 355 356 western and the eastern part of the area are different (complex terrain versus large-scale system). Overall, the results indicate that, at least for this study, as long as the terrain conditions (slope, 357 358 elevation, and orientation) are similar, the DNNs developed based on one single location can be applied with similar prediction skill for locations that are as far as 520 km (equal to more than 40 359

grid cells in the WRF output used in this study) to predict the variables assessed in this study. 360 The results also suggest that when implementing the NN-based algorithm into the WRF model, if 361 362 a number of grid cells are over a homogenous region, one may not need to train the NN over every grid cell. This will significantly save computing time because the training process takes the 363 majority of the computing resource (see below). While we show results predicted by HAC in 364 January here, we find similar conclusion from HPC prediction and both HAC and HPC 365 predictions in July, expect that the prediction skills are even better in July for the adjacent 366 locations. 367

368 3.5 DNN training and prediction time

369 Table 2 shows the number of epochs and time required for training FNN, HPC, and HAC for various numbers of training years. Because of the early stopping criterion, the number of training 370 371 epochs performed by different methods is not same. Despite setting the maximum epochs to 1000, all these methods terminate within 178 epochs. We observed that HPC performs more 372 373 training epochs than do FFN and HAC: given the same optimizer and learning rate for all the 374 methods, HPC has a better learning capability because it can improve the validation error more than HAC and FNN can. For a given set of training data, the difference in the training time per 375 epoch can be attributed to the number of trainable parameters in FNN, HPC, and HAC (10,693, 376 377 16,597, and 26,197, respectively). As we increase the size of training data, the training time per epoch increases significantly for all three DNN models. The increase also depends on the 378 number of parameters in the model. For example, increasing the training data from 1 year to 20 379 380 years increases the training time per epoch from 1.4 seconds to 11.4 seconds for FNN, from 1.1 seconds to 17.4 seconds, and from 1.4 seconds to 19.6 seconds for HPC and HAC, respectively. 381

382 The prediction times of FNN, HPC, and HAC are within 0.5 seconds for one-year data, making these models promising for PBL emulation deployment. The difference in the prediction time 383 384 between models can be attributed to the number of parameters in the DNNs: the larger the number of parameters, the longer the prediction time. For example, the prediction times for FFN 385 386 are below 0.2 seconds when using different numbers of years for training, while those for HAC 387 are around 0.4 seconds. Despite the difference in the number of training years, the number of 388 parameters for a given model is fixed. Therefore, once the model is trained, the DNN prediction 389 time depends only on the model and the number of points in the test data (1 year in this study).

Theoretically, for the given model and the test data, the prediction time should be constant even with different amounts of training dataset. However, we observed slight variations in the prediction times that range from 0.17 to 0.29 seconds for FNN, 0.30 to 0.34 seconds for HPC, and 0.36 to 0.42 seconds for HAC, which can be attributed to the system software.

394 4 Summary and Discussion

395 This study developed DNNs for emulating the YSU PBL parameterization that is used by the WRF model. Two of the DDNs take into account the domain-specific features (e.g., nonlocal 396 397 mixing in terms of vertical dependence between multiple PBL layers. The input and output data 398 for the DNNs are taken from a set of 22-year-long WRF simulations. We developed the DNNs 399 based on a midwestern location in the United States. We found that the domain-aware DNNs can reproduce the vertical profiles of wind, temperature, and water vapor mixing ratio with high 400 401 accuracy yet require fewer data than the traditional DNN, which does not care about the domain-402 specific features. The training process takes the majority of the computing time. Once trained, 403 the model can quickly predict the variables with decent accuracy. This ability makes the deep 404 neural network appealing for parameterization emulator.

Following the same architecture that we developed for Logan, Kansas, we also built DNNs for one location at Alaska. The results share the same conclusion as we have seen for the Logan site. For example, among the three DNNs, HPC and HAC show much better skill with smaller RMSEs and higher correlations than does FFN. The wind profiles are more difficult to predict than the profiles of temperature and water vapor. For FFN, the prediction accuracy increases with more training data; for HPC and HAC, the prediction skill stays similar when having six or more years of training data.

While we trained our DNNs over individual locations in this study using only one computing node (with multiple processors), there are 300,000 grid cells over our WRF model domain, which simulated the North American continent as a horizontal resolution of 12 km. To train a model for all the grid cells or all the homogeneous regions over this large domain, we will need to scale up the algorithm to hundreds if not thousands of computing nodes to accelerate the training time and the make the entire NN-based simulation faster than the original parameterization.

The ultimate goal of this project is to build an NN-based algorithm to empirically understand the 419 process in the numerical weather and climate models and to replace the PBL parameterization 420 421 and other time-consuming parameterizations that were derived from observational studies. The 422 DNNs developed in this study can provide numerically efficient solutions to a wide range of problems in environmental numerical models where lengthy, complicated calculations describing 423 424 physical processes must be repeated frequently or need a large ensemble of simulations to represent uncertainty. A possible future direction for this research is implementing these NN-425 based schemes in WRF for a new generation of hybrid regional-scale weather/climate models 426 that fully represent the physics at a very high spatial resolution at a fast computational time so as 427 to provide the means for generating large ensemble model runs. 428

Data and code availability. The data used and the code developed in this study are available at
https://github.com/pbalapra/dl-pbl.

431 *Author contributions.* JW participated in the entire project by providing domain expertise and

analyzing the results from the NN-based emulator. PB developed the deep neural networks and

433 conducted the experiments. RK proposed the idea of this project and provided high-level

434 guidance and insight for the entire study.

435 *Competing interests.* The authors declare that they have no conflict of interest.

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440 **References**

Attali, J. G., and Pagès, G.: Approximations of functions by a multilayer perception: A new
approach, Neural Networks, 6, 1069–1081, 1997.

443

Chevallier, F., Chéruy, F., Scott, N. A., and Chédin, A.: A neural network approach for a fast and
accurate computation of longwave radiative budget, J. Appl. Meteorol., 37, 1385–1397, 1998.

Chen, T., and Chen, H.: Approximation capability to functions of several variables, nonlinear
functionals and operators by radial basis function neural networks, Neural Networks, 6, 904–
910, 1995a.

450

Chen, T., and Chen, H.: Universal approximation to nonlinear operators by neural networks with
arbitrary activation function and its application to dynamical systems, Neural Networks, 6, 911–
917, 1995b.

454

Chevallier, F., Morcrette, J.-J., Chéruy, F., and Scott, N. A.: Use of a neural-network-based
longwave radiative transfer scheme in the EMCWF atmospheric model, Q. J. R. Meteorol. Soc.,
126, 761–776, 2000.

458

Collins, W., and Tissot, P.: An artificial neural network model to predict thunderstorms within
400 km² South Texas domains, Meteorol. Appl., 22 (3), 650-665, 2015.

461

Cohen, A.E., Cavallo, S.M., Coniglio, M.C., Brooks, H.E.: A review of planetary boundary layer
parameterization schemes and their sensitivity in simulating a southeast U.S. cold season severe
weather environment, Weather Forecast., 30, 591-612, 2015.

465

466 Cybenko, G.: Approximation by superposition of sigmoidal functions, Math. Control Signals
467 Syst., 2, 303–314, 1989.

468

469 Hauser, A., and Bühlmann, P.: Characterization and greedy learning of interventional Markov

470 equivalence classes of directed acyclic graphs. J. Mach. Learn. Res., 13, 2409–2464, 2002.

471

Hong, S.-Y., Noh, S.Y., and Dudhia, J.: A new vertical diffusion package with an explicit
treatment of entrainment processes. Mon. Wea. Rev., 134, 2318–2341, 2006.

474

475 Hornik, K.: Approximation capabilities of multilayer feedforward network, Neural Networks, 4,
476 251–257, 1991.

478	Jiang, GQ., Xu, J., and Wei, J.: A deep learning algorithm of neural network for the
479	parameterization of typhoon-ocean feedback in typhoon forecast models, Geophysical Research
480	Letters, 45, 3706–3716, 2018.
481	
482	Kheshgi, H. S., Jain, A. K., Kotamarthi, V. R., and Wuebbles, D. J.: Future atmospheric methane
483	concentrations in the context of the stabilization of greenhouse gas concentrations. Journal of
484	Geophysical Research: Atmospheres, 104, D16: 19183–19190, 1999.
485	
486	Krasnopolsky, V. M., Fox-Rabinovitz, M. S., and Chalikov, D. V.: New approach to calculation
487	of atmospheric model physics: Accurate and fast neural network emulation of long wave
488	radiation in a climate model, Mon. Weather Rev., 133, 1370–1383, 2005.
489	
490	Krasnopolsky, V. M., and Fox-Rabinovitz, M. S.: Complex hybrid models combining
491	deterministic and machine learning components for numerical climate modeling and weather
492	prediction, Neural Networks, 19, 122–134, 2006.
493	
494	Krasnopolsky, V. M., Fox-Rabinovitz, M.S., and Belochitski, A. A.: Using ensemble of neural
495	networks to learn stochastic convection parameterizations for climate and numerical weather
496	prediction models from data simulated by a cloud resolving model. Adv. Artif. Neural. Syst., 1-
497	13, 2013.
498	
499	Krasnopolsky, V. M., S. Nadiga, A. Mehra, E. Bayler, and D. Behringer: Neural networks
500	technique for filling gaps in satellite measurements: Application to ocean color observations,
501	Computational Intelligence and Neuroscience, 2016, Article ID 6156513, 9 pages, 2016.
502	doi:10.1155/2016/6156513
503	
504	Krasnopolsky, V. M., J. Middlecoff, J. Beck, I. Geresdi, and Z. Toth. A neural network emulator
505	for microphysics schemes. 97th AMS annual meeting, Seattle, WA. January 24, 2017.
506	

507	Lee, L. A., Carslaw, K. S., Pringle, K. J., Mann, G. M., and Spracklen, D. V.: Emulation of a
508	complex global aerosol model to quantify sensitivity to uncertain parameters, Atmos. Chem.
509	Phys., 11, 12,253–12,273, 2011.
510	
511	Leeds, W. B., Wikle, C. K., Fiechter, J., Brown, J., and Milliff, R. F.: Modeling 3D spatio-
512	temporal biogeochemical processes with a forest of 1D statistical emulators. Environmetrics,
513	24(1): 1–12, 2013.
514	
515	McFarlane, N.: Parameterizations: representing key processes in climate models without
516	resolving them. Wiley Interdisciplinary Reviews: Climate Change, 2 (4): 482–497, 2011.
517	
518	Scher, S.: Toward data-driven weather and climate forecasting: Approximating a simple general
519	circulation model with deep learning. Geophysical Research Letters, 45, 12,616–12,622, 2018.
520	
521	Thompson, G., Field, P.R., Rasmussen, R.M., Hall, W.D.: Explicit forecasts of winter
522	precipitation using an improved bulk microphysics scheme. Part II: Implementation of a new
523	snow parameterization, Mon. Weather Rev. 136, 5095–5115, 2008.
524	
525	Wang, J., and Kotamarthi, V. R.: Downscaling with a nested regional climate model in near-
526	surface fields over the contiguous United States, Journal of Geophysical Research, Atmosphere,
527	119, 8778–8797, 2014.
528	
529	Williams, P. D.: Modelling climate change: the role of unresolved processes. Philosophical
530	Transactions of the Royal Society A: Mathematical, Physical and Engineering Sciences 363

531 (1837): 2931–2946, 2005.

532 Table 1: Inputs and outputs for the NN developed in this study. The variable names of

533 these inputs and outputs in the WRF are shown in the parentheses.

Output Variable				
zonal wind (U)				
meridional wind (V)				
temperature (tk)				
water vapor mixing ratio (QVAPOR)				
Ground temp (TSK)				
Soil temperature at 2 m below ground (TSLB)				

DNN Type	Training Data (years)	Training Time (s)	Number of Epochs	Training Time (s) per Epoch	Prediction Time (s) for 1 Year (2005)
FNN	1	85.969	61	1.409	0.197
FNN	2	137.359	47	2.923	0.196
FNN	6	376.209	70	5.374	0.171
FNN	12	199.468	23	8.673	0.193
FNN	20	306.665	27	11.358	0.199
HPC HPC HPC HPC	1 2 6 12	199.152 454.225 1233.908 1225.880	178 91 133 88	1.119 4.991 9.278 13.930	0.336 0.343 0.317 0.302
HPC	20	1181.716	68	17.378	0.331
HAC HAC HAC	1 2 6	131.104 468.884 870.753	95 85 80	1.380 5.516 10.884	0.366 0.411 0.406
HAC	12	737.921	47	15.700	0.420
HAC	20	1351.898	69	19.593	0.381

Table 2: Training and prediction time (unit: seconds) for the three DNNs using different
lengths of training data. The predicted period is for one year (2005).



Figure 1: Three variants of DNN developed in this study. Red, yellow, and purple indicate 541 the input layer (16 near-surface variables), output layers, and hidden layers, respectively. 542 (a) fully connected feed forward neural network (FFN), which has only one output layer 543 with 85 variables (5 variables for each of the 17 WRF model vertical levels), and 17 hidden 544 layers which only consider the near-surface variables as inputs. (b) hierarchically 545 connected network with previous layer only connection (HPC), which has 17 output layers 546 (corresponding to the PBL levels) with each of them having 5 variables, and 17 hidden 547 layers with each them considering both near-surface variables and output variables from 548 previous output layer as inputs. (c) hierarchically connected network with all previous 549 layers connection (HAC), same as HPC, but each hidden layer also considers output 550 variables from *all* previous output layers as inputs. 551



Figure 2: Temperature and water vapor mixing ratio from the observation and three DNN
predictions: FFN, HPC, and HAC in January and July of 2005 at 3 PM and 12 AM local
time. The y-axis uses log scale. The training data are from 20 years (1984 to 2003) of 3hourly WRF output. The lower and upper dash lines show the lowest and highest (5th and
95th percentile) PBL heights at that particular time. For example, in January 12 AM, the
lowest PBL height is about 19 m, while the highest PBL height is about 365 m.



Figure 3: RMSE and correlations for time series of temperature and water vapor vertical profiles within the PBL predicted by the three DNNs compared with the observations. The vertical lines show the range of RMSEs and correlations when considering the lowest and highest PBL heights at each particular time (shown by the dashed horizontal lines in Figure 2). The training data are 3-hourly WRF output from 1984 to 2003.



566 Figure 4: Same as Figure 2 but for wind direction and wind speed.



568 Figure 5: Same as Figure 3 but for wind components.



569

Figure 6: RMSEs for temperature, water vapor, and wind components at midnight of January using three DNNs. Left y-axis is for RMSEs of HAC and HPC; right y-axis is for RMSE of FFN. The RMSEs are calculated along the time series below the PBL height for January midnight at local time. The lower and upper end of the dash lines are RMSEs that consider the lowest and highest PBL heights as shown in Figure 2.



576 Figure 7: Same as Figure 6 but for Pearson correlations.





Figure 8: Terrain height (in meters) over the tested area; and normalized RMSEs in %
(relative to their corresponding observations) of HAC predicted temperature, water vapor

mixing ratio, wind direction and speed at midnight of January. The star shows where the
DNNs are developed (Logan, Kansas).

583



Figure 9: Person correlations between observed and HAC predicted temperature, water
mixing ratio, wind direction and speed for midnight of January in 2005.