

We thank the reviewers for their insightful comments. The reviewers' comments are repeated below in italics, followed by point-by-point responses to each. We also have improved the figures and text covering sea ice after helpful comments from a collaborator (to be added as an author upon revision).

## **Response to Reviewer #1**

*This paper shows the impact of horizontal resolution on ocean-sea-ice model simulations by comparing the outputs from four models at low and high horizontal resolutions. Several biases in the low-resolutions models are improved when the resolution becomes high. The impacts of model resolution on temperature and salinity bias depend on the model. Although the model configuration differences between low and high resolutions differ among the four models, several improvements in high-resolution models compared to low-resolution models seem to be robust. On the one hand, the biases of temperature and salinity fields depend on the model and region. The report in this paper is informative for readers of GMD. I recommend this paper for publication after some minor modifications.*

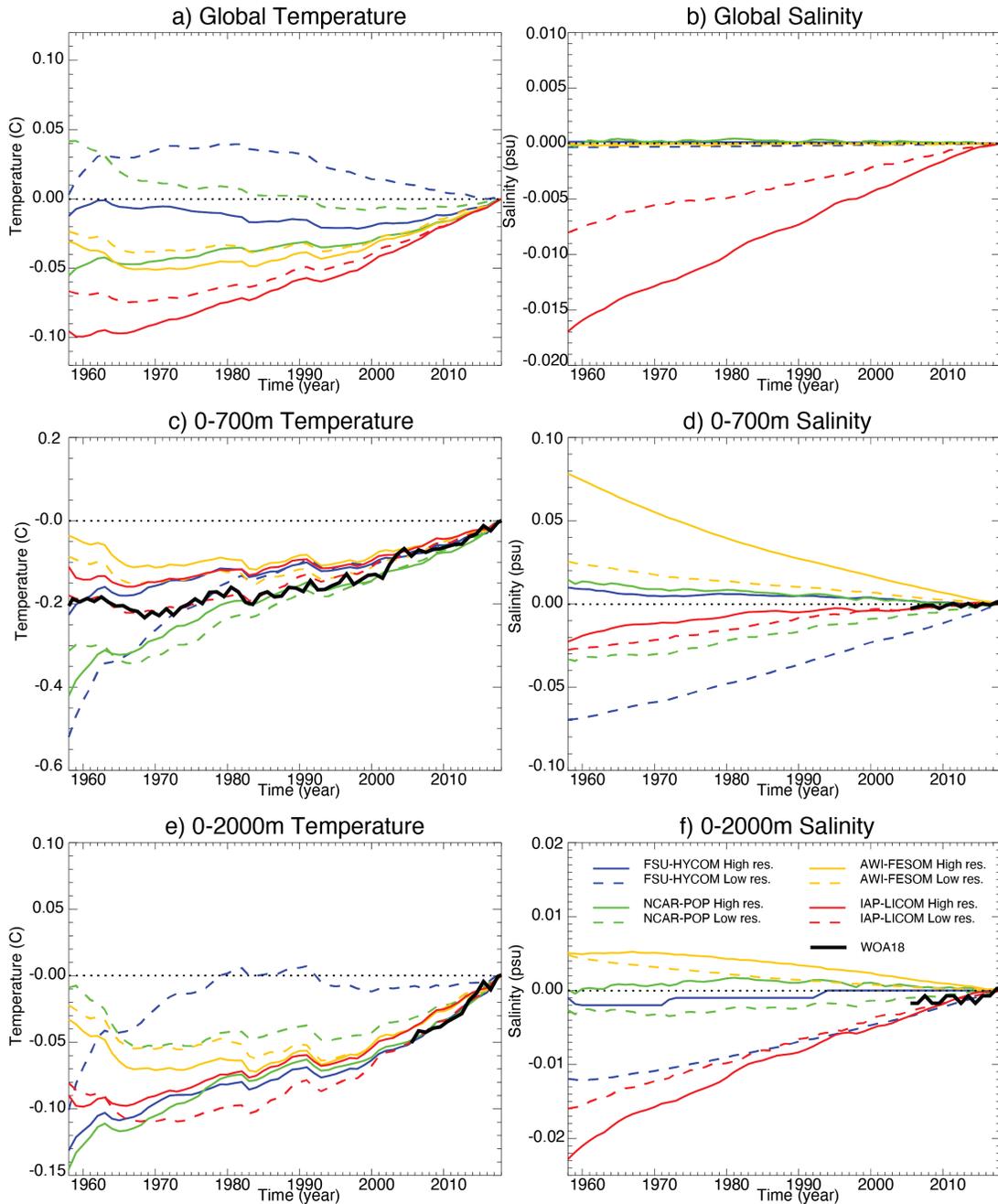
### Comments

*The time evolution of the global temperature was examined in Section 3.2. Levitus et al. (2005) showed a warming of the world ocean over 1955-2003 based on the observations. The warmings in the models seem to be consistent with the observation including its decadal variations. As well as the sea surface height, the impact of global warming should be discussed.*

A primary focus of Section 3 is on simulation drift, which tends to be much larger than forced variability in the first few forcing cycles of OMIP simulations for many fields, particularly global ocean temperature (Danabasoglu et al. 2014; Griffies et al. 2014). With only one forcing cycle available from the high-resolution OMIP-2 runs, it can be difficult to distinguish forced variability from model drift. The OMIP simulations are initialized from observed climatologies, not from observed 1958 conditions, and so it is not expected that they should accurately simulate the full-depth evolution of the ocean over the 1958-2018 forcing cycle. In the real world, this evolution reflects the cumulative effect of hundreds or thousands of years of surface forcing. Furthermore, the JRA55-do forcing dataset includes adjustments designed to yield zero net climatological heat/freshwater flux into the ocean for the 1988-2007 time period (Tsujino et al. 2018). Thus, an ocean model that perfectly replicates the observed sea surface state should exhibit zero net ocean heating over that 20-year time period when forced with JRA55-do. This is unrealistic, but it is a reasonable choice in the dataset construction given the large uncertainty in the climatological net surface forcing of the ocean (Tsujino et al. 2018). While it is problematic to assess the realism of long-term ocean warming in OMIP simulations for all the reasons outlined above, the JRA55-do forcing should reflect variations in the global energy imbalance that produce interannual to decadal changes in the rate of ocean warming (Trenberth et al. 2014).

In response to the reviewer, we now include two versions of Figure 4 (time evolution of global temperature and salinity) in the paper. The original version (showing anomalies relative to initial conditions) highlights model drift away from observed ocean climatology. The new version (showing anomalies relative to 2018) highlights the forced ocean variations of the last few decades of simulation (see below). A comparison with observed heat/salt content change from the WOA18 product in panels c-f shows that high resolution improves the fidelity and reduces the spread of forced ocean heat content change (particularly 0-2000m heat content) over recent decades. Figures 6-8 (time evolution of ocean temperature as a function of depth)

are left unchanged because they are intended to illuminate the geographical regions where model drift (i.e., bias growth) is most pronounced, not where forced ocean heat uptake occurs.

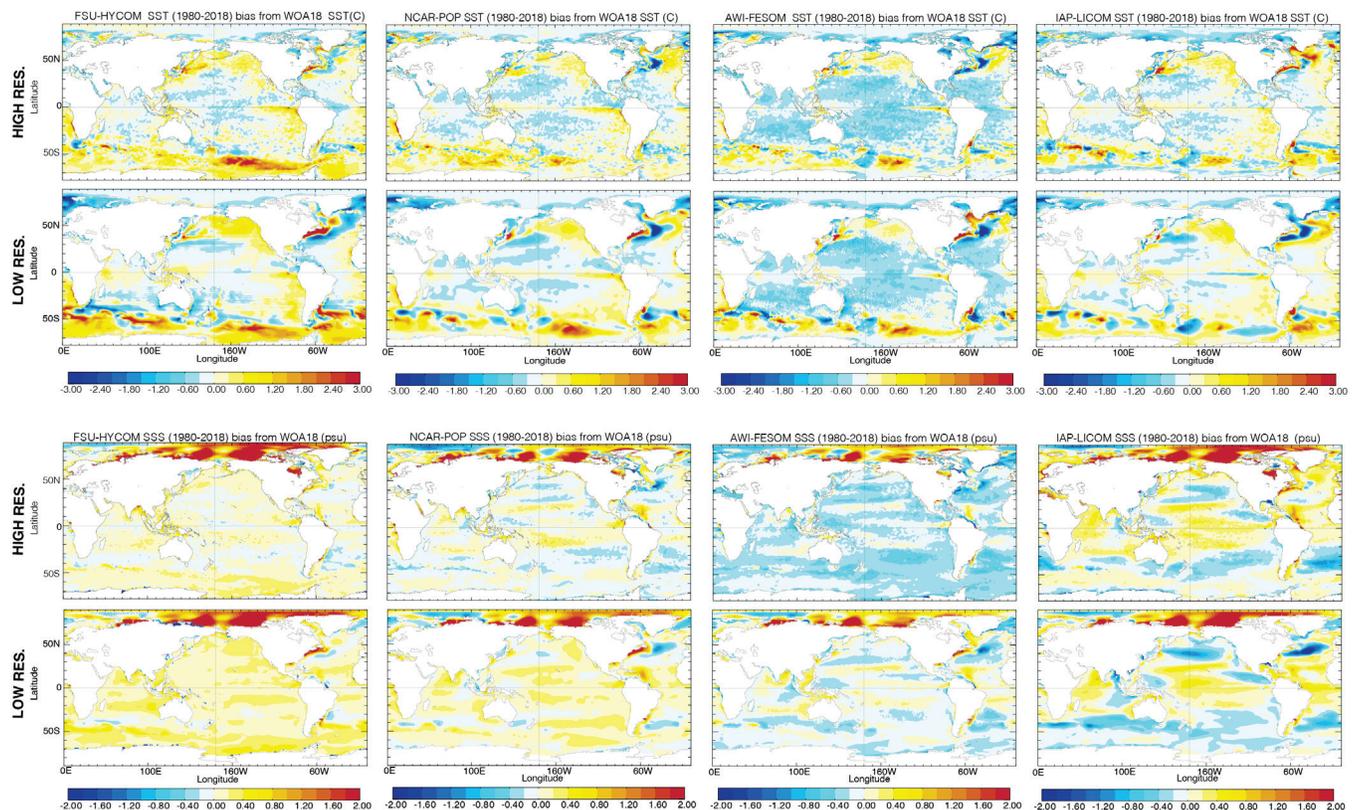


Version 2 of Figure 4, to be included in the manuscript as Figure 5

Figure 9 shows the surface temperature and salinity differences between the simulated fields over 2009-2018 and the initial fields, which discusses about how the simulated field changes from its initial state. Global warming appears to influence much the temperature difference. The model bias compared to the observation should be correctly verified. The simulated surface temperature and salinity are constrained to atmospheric forcing and the restoring to the observation respectively. Because these constrain are not

perfect. Therefore, it is worth verifying the surface temperature and salinity biases compared with recent observations such as Argo float in the same period.

The SST/SSS bias plot (Fig. 9) has been replotted using concurrent observations (WOA18) spanning 1980-2018 to match Figs. 10, 11, and 12. The new plot differs somewhat from the old, but still indicates that model bias is essentially a reflection of model drift away from the initial conditions. This bias generally sets up within a decade or two of initialization (Figs. 6-9). The discussion in the paper will be updated accordingly.



Revised Figure 9

Minor comments

There are not a few typos in the texts. The lead-author should be responsible to make readable paper. In Section 2, the model descriptions of four models are shown. It seems that different authors responsible to each model wrote the texts for each configuration, which makes the texts unreadable. The configuration of each model should be written in the same writing style.

Will do.

GMD recommend the unit with negative exponent (e.g.  $m s^{-1}$  instead of  $m/s$ ).

To be corrected.

L82: 'AMOC' appears first. 'Atlantic meridional overturning circulation' should be shown.

Will do.

*L98:  $\sigma_2$  should be defined.*

Will do.

*L107:  $\Delta x$  should be defined such as 'zonal grid spacing'.*

Will provide complete definition of  $\Delta x$ .

*L145-146: 'm2 s-1' should be modified by using superscript.*

Will do.

*L313: 'section 2b,c' should be 'section 3.2 and 3.3'.*

Will do.

*Sec.3.4: The surface temperature and salinity averaged over 2009-2018 are compared to the initial state (Fig. 9). On the one hand, the vertical structures averaged over 1980-2018 are compared to the initial state (Fig. 10, 11, 12). Why are the average periods different?*

Good point. In light of the reviewer's earlier comment on section, Fig. 9 was revised with a comparison to the 1980-2018 average period.

*Sec.3.7: Please plot the EKE time series along the ACC, if possible. Were the EKE in the models were increased by the zonal wind intensification since the 1970s.*

The difficulty with an evaluation of the EKE response to winds is that there is also an initial condition transient in the buoyancy gradients across the ACC due to the model drift discussed above (see Figure 14c illustrating the buoyancy drift). Since the MOC across the ACC is equilibrating for a significant fraction of the whole time of the high-resolution integration, it is difficult to interpret the mechanisms leading to EKE response and it is somewhat beyond the scope of this paper. Thus, we anticipate that this figure would be quite confusing to the readers. We will expand the discussion and refer to the recent literature.

*L580: I cannot find Ollitrault and Colin de Verdière, 2014 in References.*

The reference will be added.

*Fig. 27: Please add the curves of heat transport of IAP-LICOM high resolution in Fig. 27.*

Will do.

*'summer' and 'winter' in the captions of Fig. 30 and 31 should be 'winter' and 'summer'*

*respectively.*

To be corrected.

*L242, 553, 1170: Renault et al. (2019) should be Renault et al. (2020).*

To be corrected.

*L909-918: Please confirm the sequence of references.*

To be corrected.

*L963 includes two references.*

To be corrected.

*L1204 includes two references.*

To be corrected.

*Table.1: Please show the initial conditions.*

To be added.

*Table.1: 'Relative wind stress' should be 'Absolute wind stress' for FSU-HYCOM'.*

To be corrected.

## **Response to Reviewer #2**

### General Comments

*This is an important paper for the ocean and climate modelling community feeding into wider decisions on the benefits of high resolution ocean components to address climate science questions. The conclusion that the impact of resolution is consistent across model families in terms of circulation but model dependent in terms of temperature and salinity biases is an important one. My main comment is that the high and low resolution versions of the models are not always comparable due to differences in vertical as well as horizontal resolution. It would have been preferable to have changed only the horizontal resolution and associated physics so that the effects of horizontal resolution can be isolated as stated in the abstract. I would suggest that at very least, this issue should be addressed in the model description and discussion in the context of Stewart et al. (Ocean Modelling, 2017) and included as a future recommendation.*

Will do. This was an opportunistic study taking advantage of ongoing simulations rather than a dedicated parallel set of simulations. As discussed below, more often than not, the low- and high-resolution were configured independently for different scientific goals and followed different development trajectories,

when ideally they should have been designed in tandem. We will add a paragraph on the lessons learned that can be used to improve the design of future high- and low-resolution comparisons.

*I would recommend this paper for publication assuming that this and other specific issues below are addressed.*

Specific comments:

*L21: See discussion above on isolating the effects of horizontal resolution*

We will include a discussion on isolating the effects of horizontal resolution.

*L86: Title would be better as Model Descriptions*

Agreed.

*L93: A summary table of the four models at both resolutions would be useful for the reader*

The four models are summarized in Table 1. The table will be updated to include the climatologies used to initialize the models.

*L117: Why is FSU model 41 layers in low res and 36 layers in high res? (see above)*

The 36-layer high-resolution configuration was at the time our default configuration and we retained it to compare to previous runs performed with other atmospheric forcings. The 41-layer coarse-resolution runs were performed afterward for inclusion in the Tsujino et al. (2020) using the latest vertical grid with all the additional layers in the upper ocean. This will be addressed in the revised manuscript.

*L157: What is the impact in the NCAR model of two additional levels as well as partial step formulation?*

The additional (250m thick) vertical levels in high resolution POP increase the maximum ocean depth from 5500 m to 6000 m, allowing for a more realistic representation of deep ocean features resolved by the 0.1° grid. The partial bottom cell formulation improves the representation of deep bathymetric slopes. Neither has been tested in the 1° version, and so the impacts are unknown, but are presumed to be small compared to the horizontal grid refinement which includes overall improved representation of small-scale bathymetric features (even without the minor changes to vertical discretization). It should be noted that the bottom depths in the POP model description (section 2.2) were incorrect and should have been 5500m (1-degree) and 6000m (0.1-degree). This will be corrected in the revised manuscript.

*L207: I couldn't see any mention in the description of the lack of salt flux normalization in IAP?*

To be added.

*L213: What is the impact of changing from 30 to 55 levels in the high resolution?*

*It seems to me that this is more consistent with resolving baroclinic modes at high resolution (Stewart et al., 2017) but 55 levels is probably insufficient to resolve the first baroclinic mode.*

The main goal was to increase the vertical resolution in the deep ocean to improve the simulation of the deep circulation and the AMOC transport. The AMOC transport is better represented at 26.5°N in the high-resolution CAS-LICOM experiment than in the low-resolution experiment (Figure 26). We do agree with the reviewer that 55 levels may be insufficient to resolve the first baroclinic mode, but the 55 levels were dictated by the computing resources at hand. This will be discussed in the revised manuscript.

*L229: Title would be better as Temporal evolution and model drift*

Agreed.

*L249: Is there any comment on why the global average KE is so different between NCAR and IAP (which should be fairly comparable in terms of resolution)?*

This is under investigation by the LICOM team, but it most likely due to the numerics, especially the two-step shape preservation advection scheme used in the momentum equations. The Euler-backward finite difference scheme used in the advection scheme is quite diffusive and the LICOM team is exploring ways to decrease the frequency of which the Euler-backward finite difference scheme is used. In addition, CAS-LICOM uses a relatively large horizontal diffusivity. This will be discussed in the revised manuscript.

*L258: Add this discussion of the salt flux normalization to mode description and specifically in description of IAP*

Will do.

*L275: Please correct 'all high resolution models warm faster over the upper 2000m and global temperature than their lower-resolution partners'. This doesn't appear to be true for FSU.*

Will do.

*L284: I understand that including this is beyond the scope of this paper, but the authors could recommend and pass to the discussion, that the vertical heat and salt budgets are compared in future as in Griffies et al. (J. Climate, 2015) and Von Storch et al. (Ocean Modelling, 2016)*

Will do.

*L335: Is it possible to hypothesise what leads to reductions or increases in the biases and the different model behaviour?*

We can hypothesize that the thermocline bias is related to the representation of vertical eddy heat flux (Griffies et al. 2015), which tends to be stronger and more realistic in high resolution simulations (see revised Figure 4 in response to reviewer #1). The degradation in thermocline bias in POP high resolution could also be due to missing submesoscale physics, which are parameterized in low resolution, but absent

in high resolution, versions of that model. Such hypotheses will be added to the discussion in the revised manuscript.

*L417: is the similarity of the high resolution model and RAPID AMOC evolutions statistically significant or can it be shown to link with reduced deep convection? Are the convection patterns different across resolutions? I think without further discussion, this point is rather speculative.*

Agreed – the sentence will be reworded accordingly.

*L445: Should figure 14c be labelled as AABW cell?*

We prefer global MOC transport (see Lumpkin and Speer (2007) and Talley (2013) for observed values). Figure 14c will be relabeled as “global MOC transport at 34°S” and the figure caption modified to reflect that the negative values in global MOC transport are due to the northward-flowing AABW below the southward return flow.

**Figure 14:** Modeled annual mean Atlantic meridional overturning circulation (AMOC) transports at a) 26.5°N and b) 34°S; and c) global Meridional overturning circulation (MOC) transports at 34°S. The global MOC transport is negative because the northward flowing Antarctic bottom water (AABW) is below the southward return flow (see Lumpkin and Speer (2007) and Talley (2013) for observed values). Solid black line in panel a) are measurements from the RAPID array.

*L502: Since the temperature and salinity fields at depth are still evolving at the end of these experiments, it seems more likely that the ACC changes are related to buoyancy drifts. Can you comment on this?*

There is indeed an initial condition transient in the buoyancy gradients across the ACC (Figure 14c) and the MOC across the ACC is equilibrating for a significant fraction of the whole time of the high-resolution integration. We will extend the discussion to reflect this point. See also the response to reviewer #1 regarding the EKE sensitivity to decadal wind stress changes.

*L528: Also comment here about the choice of vertical levels affecting the representation of baroclinic modes (Stewart et al., 2017)*

Will do.

*L539: I would argue from figure 16, that there is at least marginal improvement in the representation of the NW corner across the models but it is difficult to see in the multi panel figure*

The representation of the NW corner in these experiments was further discussed in a companion CLIVAR Variations/Exchanges article (Figure 1 of Chassignet et al., 2020). A summary sentence will be added as well as the reference to the CLIVAR article.

*L555: Please comment on Agulhas eddies in other models*

Will do.

*L710: Can you comment on the cause of the increased northward heat transport in the Atlantic by looking at the components? Presumably, it is related to a stronger Gulf Stream transport*

Northward heat transport is usually correlated with higher AMOC transport. In the high-resolution experiments, there is also an overall reduction in subtropical Atlantic upper ocean cold bias (see Fig. 23) which is likely to have an impact (Msadek et al., 2013). This will be addressed in the revised manuscript.

Msadek, R., W.E. Johns, S.G. Yeager, G. Danabasoglu, T.L. Delworth, and A. Rosati, 2013: The Atlantic Meridional Heat Transport at 26.5°N and Its Relationship with the MOC in the RAPID Array and the GFDL and NCAR Coupled Models, *J. Climate*, 26 (12), pp. 4335-4356.

*L834: Reference should be Griffies et al. (2015) see below. Include here recommendation on diagnosing vertical heat and salt budgets in future experiments?*

Will do.

*L848: I think specific comment is merited here on the aim to maintain traceability across resolution hierarchies in terms of numerics and parameterisations as well as vertical resolution*

Agreed.

*Technical corrections:*

*L106 (and similar throughout): present as  $3 \times 10^{-5}$*

Will do.

*L414: extra 'only' in the sentence*

Thank you – to be corrected

*L505: The sentence doesn't make sense as currently written*

To be rewritten.

*References:*

*Griffies, S. M., Winton, M., Anderson, W. G., Benson, R., Delworth, T. L., Dufour, C. O., Dunne, J. P., Goddard, P., Morrison, A. K., Rosati, A., Wittenberg, A. T., Yin, J. J. and Zhang, R., 2015. Impacts on Ocean Heat from Transient Mesoscale Eddies in a Hierarchy of Climate Models, *J. Clim.*, 28, 952-977.*

*Stewart, K.D., Hogg, A. McC., Griffies, S. M., Heerdegen, A. P., Ward, M. L., Spence, P., England, M. H., 2017. Vertical resolution of baroclinic modes in global ocean models, *Ocean Modell.* 113, 50-65, doi:10.1016/j.ocemod.2017.03.012.*

*von Storch, J.-S., Haak, H., Hertwig, E., Fast, I., 2016. Vertical heat and salt fluxes due to resolved and parameterised meso-scale eddies, *Ocean Modell.*, 108, 1-19, doi:10.1016/j.ocemod.2016.10.001.*