



1 **Evaluation of CH₄MOD_{wetland} and TEM models used to**
2 **estimate global CH₄ emissions from natural wetlands**

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25 Abstract

26 Reliable models are required to estimate global wetland CH₄ emissions. This study aimed to test
27 two process-based models, CH4MOD_{wetland} and TEM, against the CH₄ flux measurements of marsh,
28 swamps, peatland and coastal wetland sites across the world; specifically, model accuracy and generality
29 were evaluated for different wetland types and in different continents, and then the global CH₄ emissions
30 from 2000 to 2010 were estimated. Both models showed similar high correlations with the observed
31 seasonal CH₄ emissions, and the regression of the observed versus computed total seasonal CH₄
32 emissions resulted in R² values of 0.78 and 0.72 by CH4MOD_{wetland} and TEM, respectively. The
33 CH4MOD_{wetland} predicted more accurately in marsh, peatland and coastal wetlands, with model
34 efficiency (EF) values of 0.22, 0.55 and 0.72, respectively; however, the model showed poor performance
35 in swamps (EF<0). The TEM model predicted better in peatland and swamp, with EF values of 0.77 and
36 0.71, respectively, but it could not accurately simulate the marsh and coastal wetland (EF<0). There was
37 a good correlation between the simulated CH₄ fluxes and the observed values on most continents.
38 However, CH4MOD_{wetland} showed no correlation with the observed values in South America and Africa.
39 TEM showed no correlation with the observations in Europe. The global CH₄ emissions for the period
40 2000–2010 were estimated to be 105.31±2.72 Tg yr⁻¹ by CH4MOD_{wetland} and 134.31±0.84 Tg yr⁻¹ by
41 TEM. Both models simulated a similar spatial distribution of CH₄ emissions across the world and among
42 continents. Marsh contributes 36%–39% to global CH₄ emissions. Lakes and rivers and swamp are the
43 second and third contributors, respectively. Other wetland types account for only approximately 20% of
44 global emissions. Based on the models' generality, if we use the more accurate model to estimate each
45 continent/wetland type, we obtain a new assessment of 116.99–124.74 Tg yr⁻¹ for the global CH₄
46 emissions for the period 2000–2010.

47 1 Introduction

48 Atmospheric methane (CH₄) is the second most prevalent human-induced greenhouse gas (GHG)
49 after carbon dioxide (CO₂). It has a 28-fold greater radiative forcing than CO₂ on a 100-year horizon
50 (Myhre et al., 2013). The radiative forcing attributed to CH₄ has been re-evaluated and was reported to
51 be almost twice as high by the Intergovernmental Panel on Climate Change (IPCC)'s 5th Assessment
52 Report (AR5) than the value reported in the 4th Assessment Report (AR4), with values of 0.97 W m⁻²



53 versus 0.48 W m^{-2} , respectively (Myhre et al., 2013). This estimate considers that the emission of CH_4
54 leads to ozone production, stratospheric water vapor and CO_2 , which can affect its own lifetime (Myhre
55 et al., 2013; Shindell et al., 2012).

56 The growth rate of the atmospheric CH_4 concentration varied in different historical periods. There
57 was an exponential increase from preindustrial times to the 1980s. The growth rate decreased after the
58 1980s and was close to zero from 1999 to 2006; then, the growth rate resumed strong growth in the period
59 of 2007–2017 (Dlugokencky et al., 2009; Dlugokencky, 2016; Nisbet et al., 2019). However, the causes
60 that drive the variations in growth rate remain unclear due to the uncertainties in estimating CH_4
61 emissions and sinks (Ghosh et al., 2015; Saunio et al., 2016; Nisbet et al., 2019; Dalsøren et al., 2016).

62 Integrated at the global scale, wetlands are the largest and most uncertain source of CH_4 emitted to
63 the atmosphere (Kirschke et al., 2013; Saunio et al., 2016). These emissions represent approximately
64 30% of the total CH_4 source (Saunio et al., 2016). Bottom-up and top-down approaches are popular
65 methods for estimating global CH_4 emissions from natural wetlands. The top-down approach is based on
66 inverse models (e.g., Bousquet et al., 2006; Fraser et al., 2013; Meirink et al., 2008; Tsuruta et al., 2017;
67 Bruhwiler et al., 2014), which determine ‘optimal’ surface fluxes that best fit atmospheric CH_4
68 observations given an atmospheric transport model including chemistry, prior estimates of fluxes, and
69 their uncertainties (Kirschke et al., 2013). The bottom-up approach uses process-based models that
70 describe the relationship between the environmental factors and the processes of CH_4 production,
71 oxidation and emission using mathematical equations (e.g., Li et al., 2010; Zhu et al., 2013; Zhang et al.,
72 2002; Zhu et al., 2014; Walter and Heimann, 2000; Tian et al., 2015; Riley et al., 2011; Meng et al., 2012;
73 Zhuang et al., 2006).

74 The top-down inversions show larger uncertainties regarding CH_4 sources than the bottom models
75 (Saunio et al., 2016). Recent studies related to the bottom-up approach used an ensemble of process-
76 based models driven by the same climate forcing to estimate the global CH_4 emissions from natural
77 wetlands. For example, The Wetland and Wetland CH_4 Inter-comparison of Models Project
78 (WETCHIMP) used ten land surface models and estimated a global CH_4 emissions of $190 \pm 76 \text{ CH}_4 \text{ yr}^{-1}$
79 for the 1993–2004 period (Melton et al., 2013). In the following year, Kirschke et al. (2013) assessed a
80 large emission range of $142\text{--}287 \text{ Tg CH}_4 \text{ yr}^{-1}$ from 1980 to 2010; Saunio et al. (2016) and Poulter et al.
81 (2017) estimated global emissions of $153\text{--}227 \text{ Tg CH}_4 \text{ yr}^{-1}$ for the 2003–2012 decade and $184 \pm 22 \text{ Tg}$



82 CH₄ yr⁻¹ for the 2000–2012 decade using ensemble process-based models (Poulter et al., 2017). Saunois
83 et al. (2016) suggested that approximately 70% of the uncertainty was due to model structures and
84 parameters.

85 Natural wetland ecosystems are greatly heterogeneous on a global scale. Wetlands vary widely by
86 continent with respect to area and type (Kingsford et al., 2016; Keddy, 2010). Some wetland types have
87 higher emissions, while some emit less CH₄; this difference is because the processes of and controls on
88 CH₄ cycling differ among wetland types (Bridgham et al., 2013). For example, sedge-dominated marsh
89 or fen often emit higher CH₄ fluxes because sedges can increase methanogenic substrates as part of their
90 plant productivity and promote CH₄ transportation through their soft aerenchyma and lacunae tissues
91 (McEwing et al., 2015; Jitka et al., 2017; Bhullar et al., 2013; Joabsson and Christensen, 2001; Kwon et
92 al., 2017; King et al., 2002; Chanton, 2005). Bog soils with anaerobic incubations emit little CH₄ due to
93 the particularly high CO₂: CH₄ ratios of the end products of anaerobic carbon (Bridgham et al., 1998;
94 Galand et al., 2010; Keller and Bridgham, 2007). Coastal wetlands with high salinity usually emit less
95 CH₄ because the sulfate in the sea water inhibits CH₄ production (Bartlett et al., 1985; Delaune et al.,
96 1983; Li et al., 2016; Poffenbarger et al., 2011).

97 Model evaluation is a core part of model development and testing (Bennett et al., 2013). Based on
98 the model evaluation, the modeler must be confident that the model will fulfill its purpose (Bennett et al.,
99 2013; Rykiel, 1996). If applying process-based models for global-scale CH₄ estimations, it is necessary
100 to evaluate its performance in different wetland types and regions. This process is also helpful for
101 confirming the source of uncertainties and improving the model. However, previous studies have always
102 focused on global assessments and have overlooked model performance in different wetland types or
103 regions, which may have induced high uncertainties (Poulter et al., 2017; Saunois et al., 2016; Kirschke
104 et al., 2013; Melton et al., 2013). CH₄MOD_{wetland} (Li et al., 2010) and TEM (Zhuang et al., 2004; Melillo
105 et al., 1993; Zhuang et al., 2007; Zhuang et al., 2013) are two established process-based models that can
106 be used to simulate regional and global wetland CH₄ emissions. Both models have been validated at
107 specific sites (Zhu et al., 2013; Li et al., 2010; Li et al., 2017). However, we do not have information on
108 the accuracy and generality of the models for different wetland types and on different continents. The
109 objectives of this study were to comprehensively evaluate the model performances of CH₄MOD_{wetland}



110 and TEM for different wetland types and on different continents and then to use the models to estimate
111 global CH₄ emissions from natural wetlands.

112 **2 Methods and Materials**

113 The performance evaluation should clearly depend on the model objectives (Bennett et al., 2013).
114 The models considered in this study are aimed at estimating the annual emissions from global wetlands.
115 Therefore, the accuracy and generality of the model in simulating annual CH₄ emissions for different
116 wetland types and continents are very important in a performance evaluation. In this study, we collected
117 CH₄ flux measurements from 30 wetlands spanning the main wetland types in the world from peer-
118 reviewed literature (Table 1). A set of statistical methods was used to comprehensively evaluate the
119 performance of CH₄MOD_{wetland} and TEM in different wetland types and on different continents. Finally,
120 we extrapolated both models to estimate the global CH₄ emissions from 2000 to 2010.

121 **2.1 Model overview**

122 **2.1.1 CH₄MOD_{wetland}**

123 The CH₄MOD_{wetland} model is a process-based biogeophysical model used to simulate the processes
124 of CH₄ production, oxidation and emission from natural wetlands (Li et al., 2010). The model was
125 established based on CH₄MOD, which is used to predict CH₄ emissions from rice paddies (Huang et al.,
126 1998; Huang et al., 1997). In CH₄MOD_{wetland}, we focused on the different supply of methanogenic
127 substrates between natural wetlands and rice paddies. The methanogenic substrates were derived from
128 the root exudates, the decomposition of plant litter and the soil organic matter. The methane production
129 rates were determined by the methanogenic substrates and the influence of environmental factors,
130 including soil temperature, soil texture and soil Eh. Additionally, we adopted the influence of salinity on
131 CH₄ production to improve the model performance for the coastal wetlands (Li et al., 2016). Inputs to
132 the CH₄MOD_{wetland} model include the daily air/soil temperature, water table depth, annual above-ground
133 net primary productivity (ANPP), soil sand fraction, soil organic matter, bulk density and soil salinity.
134 The outputs are the daily and annual CH₄ production and emissions. We used the TOPMODEL
135 hydrological model to simulate the water table depth as the inputs of CH₄MOD_{wetland} (Bohn et al., 2007;



136 Li et al., 2015; Li et al., 2019; Zhu et al., 2013; Beven and Kirkby, 1979).

137 The main parameters that must be calibrated in CH4MOD_{wetland} include the vegetation index (VI),
138 the fraction of plant-mediated transport available (T_{veg}), the fraction of CH₄ oxidized during plant-
139 mediated transport (P_{ox}), the proportion of below-ground net primary productivity ($BNPP$) to the total
140 net primary productivity (NPP) (f_r), the fraction of nonstructural component in plant litter (F_N) (Table
141 S1) and the empirical constant of the influence of salinity. The model parametrization and main
142 parameters are described in Supplementary Material S1.

143 2.1.2 TEM

144 The terrestrial ecosystem model (TEM) is another process-based biogeochemical model that
145 couples carbon, nitrogen, water, and heat processes in terrestrial ecosystems to simulate ecosystem
146 carbon and nitrogen dynamics (Melillo et al., 1993; Zhuang et al., 2007; Zhuang et al., 2013). The
147 methane dynamics module was first coupled within the TEM by Zhuang et al. (2004) to explicitly
148 simulate the process of methane production (methanogenesis), oxidation (methanotrophy) and transport
149 between the soil and the atmosphere. Methane production is assumed to occur only in saturated zones
150 and is regulated by organic substrate, soil thermal conditions, soil PH, and soil redox potentials; methane
151 oxidation, which occurs in the unsaturated zone, depends on the soil methane and oxygen
152 concentrations, temperature, moisture and redox potential. Methane transport is described by three
153 pathways in the TEM: (1) diffusion through the soil profile; (2) plant-aided transport; and (3) ebullition.
154 The TEM has also been coupled with TOPMODEL (Zhu et al., 2013). The model calibration of TEM
155 has been well documented in Supplementary Material S2 and Table S2.

156 2.2 Site information and data sources

157 2.2.1 Site information

158 We collected 30 wetland sites across the world (Table 1). The wetland sites included 6 marsh sites,
159 14 peatland sites, 6 swamp sites and 4 coastal wetland sites. Among the wetland sites, 5 sites are
160 distributed in Europe (EU), 11 sites are distributed in Asia (AS), 2 sites are distributed in Africa (AF), 3
161 sites are distributed in South America (SA) and 9 sites are distributed in North America (NA). The



162 observations were from the late 1980s to the early 2010s. For most of the wetland sites, the total amount
163 of seasonal CH₄ fluxes during the observation period was calculated by summing the daily observations.
164 The absence of CH₄ emission measurements between two adjacent days of observation was linearly
165 interpolated. For a few wetland sites, the observed seasonal CH₄ fluxes were directly obtained from the
166 literature. More details about the location, vegetation and observation periods are described in Table 1.

167 2.2.2 Wetland map

168 The global wetland distributions of different wetland types were based on the Global Lakes and
169 Wetlands Database (GLWD-3) (<http://www.wwfus.org/science/data.cfm>) (Lehner and Döll, 2004) (Fig.
170 1). According to the GLWD-3, the wetland types include : 1. lakes, 2. reservoirs, 3. rivers (we combined
171 lakes, reservoirs and rivers as a single wetland type, hereinafter referred to as lakes and rivers), 4.
172 freshwater marsh and floodplain (hereinafter referred to as marsh), 5. swamp forest and flooded forest
173 (hereinafter referred to as swamp), 6. coastal wetland, 7. saline wetland (we combined coastal wetland
174 and saline wetland as a single wetland type, hereinafter referred to as coastal wetland), 8. bog, fen and
175 mire (hereinafter referred to as peatland), 9. intermittent wetland and 10. no-specific wetland. All of the
176 observed sites (Table 1) are distributed on the wetland map (Fig. 1).

177 The global wetland area (excluding rivers) was estimated by the “Global Review of Wetland
178 Resources and Priorities for Wetland Inventory (GRoWI)” as 530-570 M ha (Spiers, 1999). We used an
179 average value, as the wetland area excluded rivers in this study. The global wetland area of rivers was
180 based on the GLWD-3. Therefore, we assumed that the global wetland area was 584 M ha, which
181 represented the wetland area for the period from 2000 to 2010. The cartography-based GLWD-3 data
182 provide a global distribution of natural wetlands at a 30-second resolution. Then, we aggregated the
183 merged map up to 0.5°×0.5° (latitude × longitude) grids. The wetland area (excluding rivers) in each
184 pixel was adjusted by the ratio of the global wetland area estimated by GRoWI and by the GLWD-3.

185 2.2.3 Driver data

186 The input climate data for the models include the daily air temperature, precipitation, cloudiness
187 and vapor pressure. The historical daily climate data were developed from the latest monthly data sets of
188 the Climatic Research Unit (CRU TS 3.10) of the University of East Anglia in the United Kingdom



189 (Harris et al., 2014).

190 The soil properties needed by the CH4MOD_{wetland} model include soil texture (percentage of soil
191 sand), bulk density, soil organic carbon content, soil temperature and soil moisture. The additional
192 information needed by the TEM model includes the percentage of soil silt and clay, soil pH and territorial
193 elevation. The soil texture data were derived from the soil map of the Food and Agriculture Organization
194 (FAO) (FAO, 2012). The soil organic carbon content and the reference bulk density of wetland soils were
195 retrieved from the Harmonized World Soil Database (HWSD) (FAO, 2008) by masking the HWSD with
196 the Global Lakes and Wetlands Database (GLWD) (Lehner and Döll, 2004). The daily soil temperature
197 data were estimated by the TEM from spatially interpolated climate data. The daily soil moisture driving
198 CH4MOD_{wetland} coupled with TOPMODEL was developed from the monthly dataset
199 (http://www.cpc.ncep.noaa.gov/soilmst/leaky_glb.htm) by temporal linear interpolation (Fan and van
200 den Dool, 2004). The soil PH was also derived from the global soil property dataset of the International
201 Geosphere-Biosphere Programme (IGBP) (Carter and Scholes, 2000).

202 The vegetation map of the IGBP was referenced to specify the vegetation parameters for
203 CH4MOD_{wetland} (Table S1) and the TEM. The map was derived from the IGBP Data and Information
204 System (DIS) DISCover Database (Belward et al., 1999; Loveland et al., 2000). The 1 km × 1 km
205 DISCover dataset was reclassified into the TEM vegetation classification scheme and then aggregated
206 into 0.5° × 0.5° grids. The annual above-ground net primary productivity used to drive CH4MOD_{wetland}
207 was from the output of the TEM model.

208 For CH4MOD_{wetland}, a high resolution of the topographic wetness index dataset (Marthews et al.,
209 2015) was used to calculate the changes in the water table. The global salinity database was from the
210 World Ocean Atlas 2009 (Antonov et al., 2010). We also used the 1 km × 1 km global elevation data
211 derived from the Shuttle Radar Topography Mission (SRTM) (Farr et al., 2007). The above data were
212 resampled to 0.5° × 0.5° grids to match the resolution of the other input data.

213 **2.3 Model evaluation**

214 We compared the observed seasonal CH₄ fluxes in the wetland sites (Table 1) and the simulated
215 seasonal CH₄ fluxes at the 0.5°×0.5° grid for the same period (described in Sect. 2.4). The statistics
216 include the determination coefficient (R²), the root mean square error (RMSE), the mean deviation



217 (RMD), the model efficiency (EF) and the coefficient of determination (CD) and were used to evaluate
218 model performance on a global scale, a continental scale and for each wetland type. Because of the site
219 limitations in Africa and South America, we combined the two continents together.

220 Two simulations with the same RMSE values may not be considered equivalent because the
221 distribution of the error among the sources may not be the same (Allen and Raktoe, 1981). We further
222 analyzed the source of the RMSE by decomposing it into three components: the bias in the modeling
223 procedure (UM), the errors due to regression (UR) and the error due to random disturbances (UE) (Allen
224 and Raktoe, 1981). The detailed description and the equations used to calculate these statistics are
225 described in Supplementary Material S3.

226 **2.4 Model extrapolation**

227 The CH4MOD_{wetland} and the TEM were used to simulate the CH₄ emissions from global wetlands at
228 a spatial resolution of 0.5°×0.5°. We established spatially explicit data for climate, soils, vegetation, land
229 use and other environmental inputs at a 0.5°×0.5° spatial resolution to facilitate the models at the global
230 scale. Both process-based models were conducted for the period of 1980-2010 in each pixel to simulate
231 the temporal spatial variations in CH₄ fluxes. In this study, we focused only on the total CH₄ emissions
232 for the period 2000-2010 because we assumed the wetland map represented the distribution of natural
233 wetlands during this period. The total CH₄ emissions from the natural wetlands, excluding the lakes and
234 rivers in each pixel, were calculated as the product of the CH₄ fluxes and the gridded wetland area. To
235 make an overall global/continental CH₄ emissions assessment, we evaluated the CH₄ emissions from
236 lakes and rivers using IPCC Tier 1 method based on the CH₄ emissions factor (IPCC, 1996) and the area
237 of lakes and rivers in each pixel.

238 We aggregated the gridded values and obtained the annual mean CH₄ emissions from each wetland
239 type and each continent by CH4MOD_{wetland} combined with IPCC Tier1 method and the TEM combined
240 with IPCC Tier1 method. Except for the two global assessments by CH4MOD_{wetland} combined with IPCC
241 Tier1 method (hereinafter referred to as “Method A”) and the TEM combined with IPCC Tier1 method
242 (hereinafter referred to as “Method B”), we made two other assessments of global CH₄ emissions by
243 choosing the more accurate model (“Method C” and “Method D”). Based on the model performance
244 evaluation (Sect. 2.3), we found a more accurate model for each wetland type and each continent. In the



245 “Method C” approach, we chose the CH₄ emissions from each continent simulated by the more accurate
246 model. In Oceania, we used the average simulated result by CH4MOD_{wetland} and the TEM because there
247 was no wetland site on this continent (Table 1). We summed the CH₄ emissions from all continents and
248 made an assessment of the global CH₄ emissions. In the “Method D” approach, we chose the CH₄
249 emissions from marsh, peatland, swamp and coastal wetlands simulated by the more accurate model. The
250 CH₄ emissions from intermittent wetlands and no-specific wetlands were used as the average result by
251 CH4MOD_{wetland} and the TEM. The CH₄ emissions from lakes and rivers were based on IPCC Tier 1
252 method. We summed the CH₄ emissions from all wetland types and made an assessment of the global
253 CH₄ emissions.

254 3. Results

255 3.1 Model evaluation

256 3.1.1 Model evaluation for global wetland sites

257 Fig. 2 shows the correlation of the modeled versus observed total amount of seasonal CH₄ emissions
258 by CH4MOD_{wetland} (Fig. 2a) and the TEM (Fig. 2b). The result indicated that the variations in the CH₄
259 emissions between sites and in different years could be delineated by both process-based models. The
260 regression of the observed versus computed total seasonal CH₄ emissions by CH4MOD_{wetland} (Fig. 2a)
261 resulted in an R² of 0.78, with a slope of 1.15 and an intercept of -0.83 g m⁻² (n=37, p<0.001). The
262 regression of the observed versus computed total seasonal CH₄ emissions by the TEM (Fig. 2b) resulted
263 in an R² of 0.66, with a slope of 0.72 and an intercept of 5.70 g m⁻² (n=37, p<0.001).

264 The statistics of the model performance of seasonal CH₄ emissions (Table 2) indicated that both
265 process-based models had the capability to simulate seasonal CH₄ emissions from natural wetlands on a
266 global scale (EF=0.58 for CH4MOD_{wetland} and EF=0.63 for the TEM). However, a discrepancy still
267 existed between the simulated and observed seasonal CH₄ emissions (RMSE=61.27% for
268 CH4MOD_{wetland} and RMSE=56.00% for the TEM). The CH4MOD_{wetland} model systematically
269 overestimated the seasonal CH₄ emissions on a global scale, with an RMD value of 11.1% (Table 2). The
270 observed seasonal amount of CH₄ emissions ranged from 0.45 to 69.00 g m⁻², with an average of 20.45
271 g m⁻². The CH4MOD_{wetland} model slightly overestimated the seasonal CH₄ emissions, with a range of



272 0.20–72.39 g m⁻² and an average value of 22.73 g m⁻² (Fig. 2a). The TEM described the average seasonal
273 CH₄ emissions on a global scale very well. The modeled amount ranged from 0.49 to 56.37 g m⁻², with
274 an average of 20.45 g m⁻², which was comparable with the observations (Fig. 2b). The RMD was
275 estimated to be 0.00%, which meant there was almost no systematic deviation between the modeled and
276 observed values (Table 2).

277 We further evaluated the model accuracy by estimating the source of the RMSE. On a global scale,
278 both models had a lower U_M (0.03 and 0.00 for CH₄MOD_{wetland} and the TEM, respectively, in Table 2),
279 which indicated that little bias was attributed to the model procedure. The TEM showed good
280 performance because the RMSE was mainly due to random disturbances (U_R=0.93). For CH₄MOD_{wetland},
281 the source of the RMSE was partly from the regression error, which was consistent with the RMD value
282 (Table 2).

283 3.1.2 Model evaluation for different continents

284 We further analyzed the model predictions by CH₄MOD_{wetland} and the TEM among different
285 continents (Fig. 3, Table 2). There was a good correlation between the simulated CH₄ fluxes and the
286 observed values on most of the continents by the two models. The R² varied between 0.21 (Fig. 3e) and
287 0.94 (Fig. 3c) for CH₄MOD_{wetland} and between 0.26 (Fig. 3d) and 0.81 (Fig. 3h) for the TEM. However,
288 there was no significant correlation between the simulated and observed values in South America and
289 Africa for CH₄MOD_{wetland} (Fig. 3g) and in Europe for the TEM (Fig. 3f). The CH₄MOD_{wetland} model
290 predicted more accurately in Asia and North America than did the TEM (Fig. 3b and 3a). The model
291 efficiency reached 0.93 in Asia and 0.48 in North America (Table 2). The TEM predicted better in North
292 America and South America and Africa than did the CH₄MOD_{wetland}, with EF values of 0.79 and 0.49,
293 respectively (Table 2). Negative values of EF were found in South America and Africa by CH₄MOD_{wetland}
294 and in Europe by the TEM (Table 2).

295 CH₄MOD_{wetland} slightly overestimated the observed emissions (RMD = 13.24%) in North America
296 and underestimated the observed emissions (RMD = -12.64%) in Asia and Europe (RMD = -29.91%)
297 (Table 2). The TEM overestimated the CH₄ emissions in South America and Africa (RMD=15.31%) and
298 underestimated the CH₄ emissions in North America (RMD=-13.03%) (Table 1). Random error was the
299 main contributor to the RMSE in Asia and Europe in CH₄MOD_{wetland} and in Asia and North America in



300 the TEM (Table 2). However, the regression error contributed most to the RMSE in North America in
301 CH4MOD_{wetland} and in South America and Africa in the TEM (Table 2).

302 3.1.3 Model evaluation for different wetland types

303 Marsh, swamp, peatland and coastal wetlands are the main types of natural wetlands. Although the
304 process-based models showed good performance in simulating the seasonal CH₄ emissions at a global
305 scale, these results do not certify that the models have good performance for each wetland type. Fig. 4
306 shows the regressions of the simulated against the observed total amount of seasonal CH₄ emissions from
307 the different wetland types. Regression analysis indicated that both models showed good performance in
308 modeling seasonal CH₄ emissions from the peatland sites (Fig. 4c and d). Most of the data points were
309 near the 1:1 line (Fig. 4c and d). Both models could explain approximately 80% of the variability in the
310 seasonal CH₄ emissions ($R^2=0.84$ and 0.77 for CH4MOD_{wetland} and the TEM, respectively) (Fig. 4c and
311 4d). The TEM showed a better model efficiency ($EF = 0.77$) than the CH4MOD_{wetland} ($EF = 0.55$), and
312 the TEM had a lower RMSE and RMD than the CH4MOD_{wetland} (Table 2) for peatland. For the other
313 wetland types, CH4MOD_{wetland} showed good performance in simulating the seasonal CH₄ emissions from
314 coastal wetlands ($EF = 0.72$), followed by marsh ($EF = 0.22$) (Table 2). However, the model could not
315 simulate the seasonal CH₄ emissions from swamps ($EF = -0.08$) (Table 2). Figs. 4a, e and d indicated the
316 same conclusion, i.e., there was no significant correlation ($p>0.05$) between the simulated and observed
317 seasonal CH₄ emissions from swamp sites (Fig. 4e). The TEM showed poor performance for the marsh
318 sites ($EF = -0.42$) and coastal wetlands ($EF = -2.26$) (Table 2); however, it showed good performance for
319 the swamp sites ($EF = 0.71$). There was no significant correlation ($p>0.05$) between the modeled and
320 observed seasonal CH₄ emissions from the marsh sites (Fig. 4b) and coastal wetland sites (Fig. 4h). In
321 conclusion, CH4MOD_{wetland} performed better for marsh and coastal wetland, while the TEM performed
322 better for peatland and swamp.

323 The sources of the RMSE varied between different wetland types during the simulation (Table 2).
324 The CH4MOD_{wetland} model had the ability to simulate CH₄ emissions from marsh, peatland and coastal
325 wetland ($EF>0$), with RMSE values of 29.44%, 77.26% and 55.46%, respectively (Table 2). For marsh
326 and peatland, the RMSE was mainly due to the regression error (Table 2). For coastal wetlands, the
327 model error contributed 24%, the regression error contributed 30%, and the random error contributed 47%



328 to the RMSE (Table 2). The TEM performed well in peatland and swamps ($EF > 0$ in Table 2). The RMSE
329 was mainly due to the random error (Table 2).

330 **3.2 Global CH₄ emissions from natural wetlands**

331 **3.2.1 Spatial pattern of global CH₄ emissions**

332 **Figs. 5a and 5b** show the distribution of the annual mean CH₄ fluxes simulated by CH4MOD_{wetland}
333 and the TEM, respectively. The models yielded similar distribution patterns of CH₄ fluxes (Fig. 5a and
334 5b). The simulated latitudinal contributions of CH₄ fluxes were consistent between the two models (Fig.
335 5a and 5b). Large fluxes were modeled in tropical regions. CH4MOD_{wetland} simulated a peak flux of 30.18
336 g m⁻² yr⁻¹ in the 10°S–0° latitudinal band, followed by fluxes over 20 g m⁻² yr⁻¹ in the 20°S–10°S
337 latitudinal band and 0°–20°N latitudinal band (Fig. 5a). A peak flux of 30.61 g m⁻² yr⁻¹ was simulated in
338 the 0°–10°N latitudinal band, followed by fluxes over 20 g m⁻² yr⁻¹ in the 20°S–0° latitudinal band and
339 10°N–20°N latitudinal band (Fig. 5b). Lower fluxes under 15 g m⁻² yr⁻¹ were modeled in the 40°N–80°N
340 latitudinal band by CH4MOD_{wetland} and in the 50°N–80°N latitudinal band by the TEM (Fig. 5a and 5b).
341 The simulation of meridional annual mean CH₄ fluxes showed the largest peak at approximately 60°W–
342 80°W and a secondary large peak at approximately 20°E–30°E (Fig. 5a and 5b).

343 The distribution of the simulated annual mean CH₄ emissions for the period 2000–2010 also showed
344 similar patterns in CH4MOD_{wetland} and the TEM. Large emissions were found in South America, South
345 Africa and the border of Canada and the United States (Fig. 5c and 5d). The latitudinal sums of CH₄
346 emissions indicated that the strongest contribution came from the tropical zone (Fig. 5c and 5d). The
347 latitudinal band of 10°S–0° contributed 22.77 Tg yr⁻¹ and 23.23 Tg yr⁻¹ CH₄ in CH4MOD_{wetland} and the
348 TEM, which accounted for 22% and 18% of the global emissions, respectively. A secondary large peak
349 was simulated in the 40°N–50°N latitudinal band, with a value of 14.64 Tg yr⁻¹ and 16.66 Tg yr⁻¹ CH₄
350 according to CH4MOD_{wetland} and the TEM, respectively. Generally, both models simulated a common
351 decline in CH₄ emissions from lower latitudes to higher latitudes (Fig. 5c and 5d). The largest peak in
352 CH₄ emissions was modeled in the 60°W–50°W meridional band, with values of 11.63 Tg yr⁻¹ in
353 CH4MOD_{wetland} (Fig. 5c) and 13.83 Tg yr⁻¹ in the TEM (Fig. 5d). This peak corresponded to the
354 longitudes of the Amazon in South America. Both models simulated secondary peaks in the 30°E–40°E



355 meridional band (Fig. 5c and 5d), which corresponded to the longitudes of Congo in Africa.

356 **3.2.2 CH₄ emissions from different continents and wetland types**

357 Table 3 provides an overview of the CH₄ emissions from different continents and wetland types
358 simulated by CH4MOD_{wetland} and the TEM. A comparison of simulated CH₄ fluxes from different
359 continents by CH4MOD_{wetland} and the TEM showed that the three highest fluxes were modeled in South
360 America, Africa and Asia (Table 3). The CH₄ fluxes simulated by CH4MOD_{wetland} were 25.8%, 18.5%
361 and 9.7% lower than those simulated by the TEM in South America, Africa and Asia, respectively. The
362 TEM simulated higher CH₄ fluxes in Europe than in North America, but the CH4MOD_{wetland} simulations
363 showed the opposite. The modeled CH₄ fluxes from European wetlands by CH4MOD_{wetland} were only
364 half of those simulated by the TEM (Table 3). For Oceania, the two models simulated similar fluxes.

365 Both models simulated the same sequence of CH₄ fluxes, which was swamp, marsh, intermittent
366 wetland, no-specific wetland, coastal wetland, and peatland (Table 3). The simulated annual mean CH₄
367 fluxes from intermittent wetlands were almost equivalent in both models. For other wetland types, the
368 TEM simulated higher CH₄ fluxes than the CH4MOD_{wetland} model (Table 3). The annual mean CH₄ flux
369 modeled by CH4MOD_{wetland} was only 20% of that modeled by the TEM (Table 3). CH4MOD_{wetland} also
370 simulated values that were 30.7%, 27.7%, 25.1% and 18.8% less than those simulated by the TEM for
371 the annual mean CH₄ fluxes from swamps, marsh, no-specific wetland and coastal wetland, respectively
372 (Table 3).

373 The global CH₄ emissions simulated by the TEM were 29 Tg yr⁻¹ higher than those simulated by
374 CH4MOD_{wetland} (Table 3). This difference depended on the differences in the CH₄ fluxes and on the
375 wetland area. The simulated results showed that half of this difference was attributed to the marsh. South
376 America contributed 30% to this difference because the simulated CH₄ fluxes differed greatly between
377 the TEM and CH4MOD_{wetland} (Table 3).

378 The two models simulated similar spatial distributions of the CH₄ emissions among different
379 wetland types and continents (Table 3). Marshes emit higher CH₄ fluxes and had the largest area. Thus,
380 marsh was the first contributor to global CH₄ emissions and contributed 36%–39% to the global CH₄
381 emissions (Table 3). Lakes and rivers as well as swamp were the second and third contributors,
382 respectively (Table 3). The CH₄ emissions from peatland, coastal wetlands, intermittent wetlands and no-



383 specific wetlands accounted for only approximately 20% of the global emissions (Table 3).

384 Although North America accounted for 36% of the global wetland area, it contributed only 22%–
385 23% to global emissions (Table 3). In contrast, the wetland area in South America accounted for 15% of
386 the global area and contributed 25%–26% to global CH₄ emissions. Asia and Africa also accounted for
387 approximately 20% of global emissions. The lowest area and emissions were found in Oceania (Table
388 3).

389 3.2.3 Global CH₄ estimations

390 The global CH₄ emissions for the period 2000–2010 were estimated to be 105.31 ± 2.72 Tg yr⁻¹ by
391 “Method A” and 134.31 ± 0.84 Tg yr⁻¹ by “Method B”. Based on the evaluation of model performance
392 (Table 2), CH4MOD_{wetland} predicted more accurately in Asia and Europe, and the TEM predicted more
393 accurately in North America and South America and Africa. Using this combination, the global CH₄
394 emissions were estimated to be 124.74 ± 1.22 Tg by “Method C”. Similarly, in “Method D”,
395 CH4MOD_{wetland} was used for simulations in marsh and coastal wetlands, and the TEM was used for
396 simulations in peatland and swamp; as a result, the global wetland CH₄ emissions were estimated to be
397 116.99 ± 2.23 Tg.

398 4. Discussion

399 4.1 Generality of CH4MOD_{wetland} and TEM

400 A lack of correspondence between the model output and observations could be partly due to the
401 observed flux data, e.g., the inevitable gap-filling of missing data points to determine annual totals
402 (Kramer et al., 2002). The results showed differences between the observed and simulated annual CH₄
403 fluxes by both CH4MOD_{wetland} and the TEM on a global scale (Fig. 2) and continent scale (Fig. 3) and
404 for different wetland types (Fig. 4). The reliability of the observed flux data was not under discussion in
405 this study. We evaluated only the model accuracy and generality across different wetland types and
406 continents.

407 On a global scale, both models fulfilled the criteria of sufficient accuracy for the ability to predict
408 CH₄ fluxes (Table 2). However, this fuzzy analysis may miss some real model performance. For the



409 model generality in different continents, CH4MOD_{wetland} had the best performance in Asia, followed by
410 North America and Europe. It had poor performance in South America and Africa, which was where
411 swamps were mainly distributed (Table 2). The TEM performed best in North America, followed by
412 South America and Africa and Asia. It had poor performance in Europe (Table 2). Each continent had its
413 main wetland types; thus, the model generality in different continents depended on its generality for
414 different types. CH4MOD_{wetland} was suitable for marsh, peatland and coastal wetlands, but it could not
415 be applied in swamps (Table 2). This limitation may be because in CH4MOD_{wetland}, only a semiempirical
416 logistic model is used to simulate plant growth (Li et al., 2010). This characteristic may induce large
417 uncertainties in simulating the growth of forests in swamps (Table 1). However, the TEM used the carbon
418 nitrogen dynamics module (CNDM) to describe the effects of photosynthesis, respiration, decomposition
419 and nutrient cycling on net primary productivity (NPP) (Melillo et al., 1993). Compared with
420 CH4MOD_{wetland}, the TEM had good performance in simulating NPP in various vegetation types (Melillo
421 et al., 1993). According to the model evaluation, the TEM was suitable for swamps and peatland but had
422 large uncertainties in marsh and coastal wetlands (Table 2). This pattern may be because the TEM focused
423 on two major wetland types: boreal tundra and forest wetland (Zhuang et al., 2004). The biochemical
424 processes in the TEM model may be suitable for peatland (tundra) and swamp (forest wetland) but not
425 suitable for marshes. For coastal wetlands, the TEM did not consider the inhibition of salinity on CH₄
426 production (Poffenbarger et al., 2011; Bartlett et al., 1987) and greatly overestimated the CH₄ fluxes
427 (Table 2). CH4MOD_{wetland} introduced the influence of salinity on CH₄ production and had good
428 performance for coastal wetlands (Table 2).

429 **4.2 Reducing uncertainties in global estimations**

430 The estimations of the global wetland CH₄ emissions had large ranges in previous studies (Zhu et
431 al., 2015). The estimations by process-based models ranged from 92 Tg yr⁻¹ (Cao et al., 1996) to 297 Tg
432 yr⁻¹ (Gedney et al., 2004) during the period of 1980–2012. Recently, an ensemble of process-based
433 models driven by the same climatic data has commonly been used to estimate global wetland CH₄
434 emissions (Melton et al., 2013; Kirschke et al., 2013; Poulter et al., 2017; Saunio et al., 2016). However,
435 the uncertainties in the model mean estimation range from 12% (Poulter et al., 2017) to 40% (Melton et
436 al., 2013). The uncertainty mainly comes from the wetland distribution and model structure and



437 parameters (Saunois et al., 2016). Estimating accurate wetland extent and its seasonal and annual
438 variations is a major challenge in present studies. The global estimations of wetland area ranged from
439 4.3 M ha to 12.9 M ha during the period of 1990 to 2005 (Melton et al., 2013). The wetland extent of 9.2
440 M ha from the GLWD excluded water bodies, and this value was ~40% higher than the wetland area
441 used in this study. That is, this difference was the main reason for the lower global estimations determined
442 in this study than those reported in previous works (Zhu et al., 2015; Melton et al., 2013; Poulter et al.,
443 2017; Saunois et al., 2016). Improving the accuracy of wetland extent and temporal variations is
444 important in reducing uncertainties in global wetland CH₄ estimations.

445 In addition to wetland area, the model structure and parameters accounted for ~70% of the total
446 uncertainties (Saunois et al., 2016). The results of the accuracy analysis showed that for CH₄MOD_{wetland},
447 regression bias accounted for 66% of the RMSE in peatland, and modeling bias accounted for 39% of
448 the RMSE in swamp; for the TEM, modeling bias and regression bias accounted for 29% and 42%,
449 respectively, of the RMSE in coastal wetland (Table 2). This result indicated that there were still
450 uncertainties in the modeling procedure, e.g., in the model mechanism or in parameterization (Zhang et
451 al., 2017; Allen and Raktoc, 1981). In the existing process-based models, which are not limited to
452 CH₄MOD_{wetland} and TEM, some important procedures should be focused on to reduce the bias due to the
453 model mechanism. For example, the mechanism of the freeze-thaw cycle is important in process-based
454 models (Wei and Wang, 2017) because of the large contribution of the CH₄ release during the nongrowing
455 season in some frozen regions (Friborg et al., 1997; Huttunen et al., 2003; Mastepanov et al., 2008; Zona
456 et al., 2016). In addition, quantifying CH₄ ebullition is important but difficult, which arises from
457 uncertainty in estimates of CH₄ emissions from peatlands (Stanley et al., 2019). Moreover, although the
458 importance of plants in CH₄ biogeochemical processes has been reported in many studies, better
459 modeling and characterization of plant community structure is needed (Bridgman et al., 2013). Finally,
460 most of the present process-based models do not have the ability to simulate CH₄ exchange from water
461 bodies, such as lakes, rivers and reservoirs, although they contribute significantly to the global budget
462 (Deemer et al., 2016). The use of the IPCC Tier method inevitably induced large uncertainties in the
463 global estimates. The above mechanisms should be incorporated into existing process-based models to
464 reduce the uncertainties in the current assessment.

465 The observational data related to processes of and controls on CH₄ production, consumption, and



466 transport also limit the model calibration and validation. The flux data of 30 wetland sites used for model
467 performance in this study are quite limited and do not represent all climatic, soil, hydrologic and
468 vegetation conditions across global natural wetlands (Table 1). During recent years, eddy covariance
469 method (Aubinet et al., 2012) is popular to observe CH₄ emissions from natural wetlands. The
470 observations in this study were used chamber method, which is widely used for CH₄ observations before
471 2010 (Table 1). There are differences in measuring CH₄ fluxes between the two methods (Chaichana et
472 al., 2018). Eddy covariance method may underestimate the fluxes (Twine et al., 2000; Sachs et al., 2010),
473 while chamber method may overestimate the fluxes (Werle and Kormann, 2001). We didn't use eddy
474 covariance data to validate the model in this study, partly because we used the chamber measurements
475 to calibrate the model. Eddy covariance data should be used for model calibration and validation in future
476 work, since it can give more accurate temporal resolution observed dataset. Furthermore, both process-
477 based models were evaluated on an annual basis rather than on a daily scale. The validation of seasonal
478 variation was not performed in this study, partly because we cannot obtain the daily step data for some
479 of the sites. Fine temporal validation against more flux datasets, especially fluxes by eddy covariance
480 experiments, and intermediate variables that control the CH₄ process are necessary in future studies (Wei
481 and Wang, 2017).

482 5. Conclusion

483 Two process-based models, CH₄MOD_{wetland} and TEM, were used to simulate annual CH₄ emissions
484 from different wetland types and continents, and their performances were evaluated. Model validation
485 showed that both models could simulate variations between different wetland sites and years. The
486 statistical analysis of model performance showed that CH₄MOD_{wetland} was capable of simulating CH₄
487 emissions from marsh, peatland and coastal wetlands, while the TEM was capable of simulating CH₄
488 emissions from peatland and swamps (model efficiency > 0). CH₄MOD_{wetland} had good performance in
489 Asia, Europe and North America, while the TEM had good performance in North America, Asia, South
490 America and Africa. The models were then used to make estimations of global wetland CH₄ emissions.
491 The CH₄ simulations of both models had good agreement in terms of the latitudinal and meridional bands.
492 The global CH₄ emissions for the period 2000–2010 were estimated to be 105.31 ± 2.72 Tg yr⁻¹ by
493 CH₄MOD_{wetland} and 134.31 ± 0.84 Tg yr⁻¹ by the TEM. If we used a more accurate model to estimate



494 each continent/wetland type based on the models' generality, the estimated global CH₄ emissions would
495 be 116.99–124.74 Tg yr⁻¹ for the period 2000–2010. The uncertainty of global wetland CH₄ assessments
496 by the process-based model approach comes from the inaccuracy of the wetland mapping area, the
497 modeling procedure and the observational limitations. Future research on accurately mapping wetland,
498 improving model mechanism and parametrization and using more observations to evaluate model
499 performance would improve global estimations.

500 **Code and data availability.**

501 The TEM and CH4MOD_{wetland} model code and model data sets (input data and model results) are
502 available on the website <https://zenodo.org/record/3594988#.XglabvkzY2w>.

503 **Author Contribution**

504 L.F. Yu pondered the rationale of the method. T. T. Li and YY. Lu developed the performed the
505 simulations. W.J. Sun, Q. Zhang, W. Zhang, GC. Wang, L.J. Yu and R. Zhang made the data collection
506 and processing. T.T. Li prepared the manuscript with contributions from all Coauthors.
507

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511 **Competing Interests**

512 The authors declare no competing interests.

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867 **Table 1. Description of observation wetland sites**

Wetland Name, Continent	Location	Wetland Type	Plant Species	Observation Period	Reference
Northeast Siberia, Russia, EU	72°22'N, 126°28'E	Peatland ^a	<i>Carex</i> spp., <i>Limprichtia revolvens</i> , <i>Meesia longisetata</i>	1999.5-1999.9	Wagner et al., 2003
Northeast Siberia, Russia, EU	71°30'N, 130°00'E	Peatland ^a	<i>Eriophorum</i> , <i>Carex</i> spp., <i>Sphagnum</i> spp., <i>Salix</i> spp.	1993.7-1993.8	Nakano et al., 2000
Northeast Siberia, Russia, EU	68°30'N, 161°24'E	Peatland ^a	<i>Larix</i> , <i>Alnus</i> spp., <i>Betula</i> spp., <i>Salix</i> spp.	1995.7-1995.8	Nakano et al., 2000
Zackenbergl, Greenland, EU	74°30'N, 21°00'W	Peatland	<i>Cassiope tetragona</i> , <i>Salix arctica</i>	1996.6-1996.8 1999.7-1999.8 2000.7-2000.8	Christensen et al., 2000; Jobsson and Christensen, 2001
Abisko, Sweden, EU	68°22'N, 19°03'E	Peatland ^a	<i>Eriophorum angustifolium</i> , <i>Carex</i> spp.	1974.6-1974.9	Svensson and Rosswall, 1984
Sanjiang Plain, China, AS	47°35'N, 133°31'E	Marsh	<i>Carex lasiocarpa</i> , <i>Deutzia angustifolia</i>	2002.6-2005.11	Hao, 2006; Song et al., 2008
Ruoergai Plateau, China, AS	32°47'N, 102°32'E	Peatland	<i>Carex multensis</i> , <i>Carex meyeriana</i>	2001.4-2001.10	Wang et al., 2002
Wuliangsu Lake, China, AS ^{&}	40°47'-41°03' N, 108°43'-108°57' E	Marsh	<i>Phragmites australis</i>	2003. 4-2003.10	Duan et al., 2005
Haiabei alpine marsh, China, AS	37°29'N, 101°12'E	Marsh	<i>Carex allivirens</i>	2002.7-2002.9	Hirota et al., 2004
Zhalong Wetland, China, AS	46°52'N-47°32'N, 123°47'E-124°37'E	Marsh	<i>Phragmites australis</i>	2009.5-2009.10	Huang et al., 2011
Liao River delta, China, AS	40°40'-41°25'N, 121°35'- 122°55'E	Coastal wetland ^b	<i>Phragmites australis</i>	1997.4-1997.11	Huang et al., 2005
Chongming Island, China, AS ^{&}	31°15'N, 121°30'E	Coastal wetland ^b	<i>Scirpus</i>	2004. 5-2004.12 2011.2-2011.12	Li et al., 2014
Guangzhou, China, AS [*]	23°01'N, 113°46'E	Coastal wetland ^c	<i>Aegiceras corniculatum</i> etc.	2005.3-2005.12	Kang et al., 2008
Haikou, China, AS [*]	19°51' N, 110°24' E	Coastal wetland ^c	<i>Bruguiera sexangula</i>	1996.1-1997.12	Ye et al., 2000
Sarawak, Malaysia, AS ^{&}	2°49'N, 111°51'E	Swamp	Flooded forest [§]	2002.8-2003.7	Melling et al., 2005
Kalimantan, Indonesia, AS	2°20'S, 113°55'E	Swamp	<i>Shorea balangeran</i>	1994.9-1995.9	Page et al., 1999; Jauchainen et al., 2005
Congo River basin, Congo, AF [*]	4°00'S-0°00', 14°00'- 18°00'E	Swamp	Flooded forest [§]	1988	Tathy et al., 1992
Congo River basin, Congo, AF [*]	0°00'-4°00'N, 14°00'- 18°00'E	Swamp	Flooded forest [§]	1988	Tathy et al., 1992
Pantanal, Brazil, SA	19°30'S, 57°00'W	Marsh	<i>Paspalum repens</i>	1998.1-1998.12	Alvalá and Kirehloff, 2000; Melack et al., 2004
Lago Calado, Brazil, SA [*]	3°15'S, 60°34'W	Swamp	Flooded forest [§]	1985	Crill et al., 1988
Central Brazilian Amazon, SA [*]	5°00'S-0°00', 50°00'-70°00'W	Swamp	Flooded forest [§]	1985	Devol et al., 1988
Alaska bethele, USA, NA	60°45'N, 161°45'W	Peatland ^a	<i>Empetrum nigrum</i> , <i>Carex aquatilis</i> , <i>Sphagnum</i> spp.	1988.7-1988.8	Bartlett et al., 1992



Alaska Prudhoe Bay, USA, NA*	70°30'N, 149°00'W	Peatland ^a	<i>Sphagnum</i> spp.	1984 1987.6-1987.10 1988.6-1988.10 1989.6-1989.10	Sebacher et al., 1986
Alaska arboretum, USA, NA	64°52'N, 147°51'W	Peatland ^a	<i>Eriophorum vaginatum</i> , <i>Carex</i> spp., <i>Sphagnum</i> spp.	1987.6-1987.10 1988.6-1988.10 1989.6-1989.10	Whalen and Reeburgh, 1992
Saskatchewan, Canada, NA	53°57'N, 105°57'W	Peatland	<i>Buckbean-Carex</i> spp.	1994.5-1994.9	Suyker et al., 1996
Michigan, USA, NA	42°27'N, 84°01'W	Peatland	<i>Scheuchzeria palustris</i> , <i>Carex oligosperma</i>	1991.1-1993.12	Shannon et al., 1996
Tootlik Lake, USA, NA	68°38'N, 149°38'W	Peatland ^a	<i>Eriophorum</i> , <i>Carex</i> spp.	1990.6-1990.8	Christensen, 1993
Hudson Bay, Canada, NA	51°18'-51°31'N, 80°28'-80°38'W	Peatland	<i>Larch</i> , <i>Black spruce</i> , <i>Sphagnum</i> spp.	1990.6-1990.10	Moore et al., 1994
Quebec, Canada, NA	54°48'N, 66°49'W	Peatland	<i>Carex</i> spp.	1989.6-1989.9	Moore et al., 1990
Mississippi, USA, NA	34°24'N, 89°50'W	Marsh	<i>Carex hyalinolepis</i> , <i>Hydrocotyle umbellata</i> , <i>Festuca obtusa</i>	2005.5-2006.7	Koh et al., 2009

* We used the average yearly CH₄ flux of the experimental year from the literature.

^a Wetland sites used for calibration.

^b These swamp sites do not have plant species information in the literature.

^c Tundra.

^d Tidal marsh.

^e Mangrove.

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Table 2. Model performance for CH4MOD_{wetland} and the TEM for different continents and wetland types

Wetland type/Continent	CH4MOD _{wetland}										TEM									
	R ²	RMSE	RMD	EF	CD	U _M	U _K	U _E	R ²	RMSE	RMD	EF	CD	U _M	U _K	U _E	n			
North America	0.80	65.82	13.24	0.48	0.46	0.04	0.60	0.36	0.80	41.93	-13.03	0.79	1.45	0.10	0.04	0.87	14			
Asia	0.94	55.79	-12.64	0.93	0.96	0.28	0.02	0.70	0.26	72.56	1.71	0.32	1.93	0.00	0.03	0.97	11			
Europe	0.21	63.66	-29.91	0.07	1.13	0.22	0.06	0.71	NS	155.88	2.42	-4.55	0.38	0.00	0.88	0.12	7			
South America & Africa	NS	59.42	42.44	-0.92	0.74	0.51	0.03	0.46	0.81	30.55	15.31	0.49	3.15	0.25	0.48	0.37	5			
Marsh	0.75	29.44	0.52	0.22	0.37	0.00	0.73	0.27	NS	39.76	-18.77	-0.42	0.95	0.22	0.17	0.61	8			
Peatland	0.84	77.26	7.60	0.55	0.47	0.01	0.66	0.33	NS	55.20	-1.64	0.77	1.09	0.01	0.02	0.97	20			
Swamp	0.32	75.22	46.68	-0.08	0.50	0.39	0.14	0.47	0.70	38.83	16.77	0.71	0.93	0.19	0.02	0.79	6			
Coastal wetland	0.80	55.46	-26.97	0.72	2.09	0.24	0.30	0.47	0.77	188.26	101.00	-2.26	0.35	0.29	0.42	0.29	4			
Global	0.78	61.27	11.13	0.58	0.58	0.03	0.43	0.54	0.66	56.00	-0.00	0.63	1.26	0.00	0.02	0.98	37			

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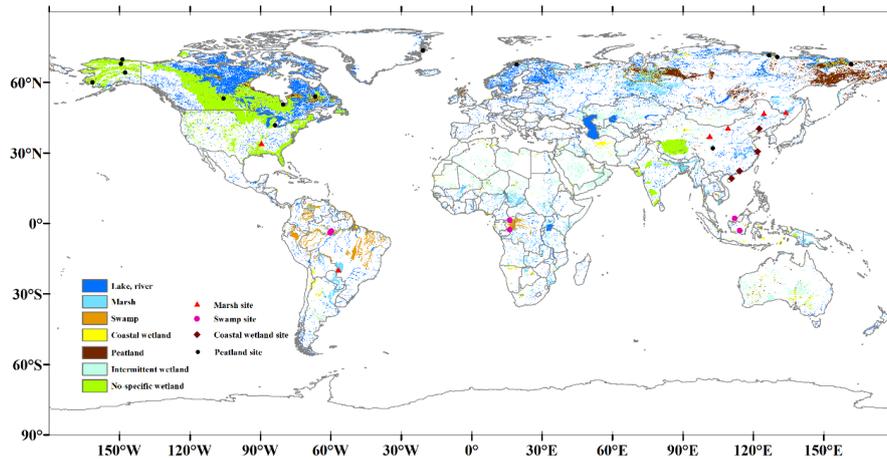
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Table 3. CH₄ simulations by CH₄MOD_{wetland} and the TEM for different continents and wetland types. All units are Tg CH₄ yr⁻¹ ± 1σ, where the standard deviation represents the interannual variation in the model estimates.

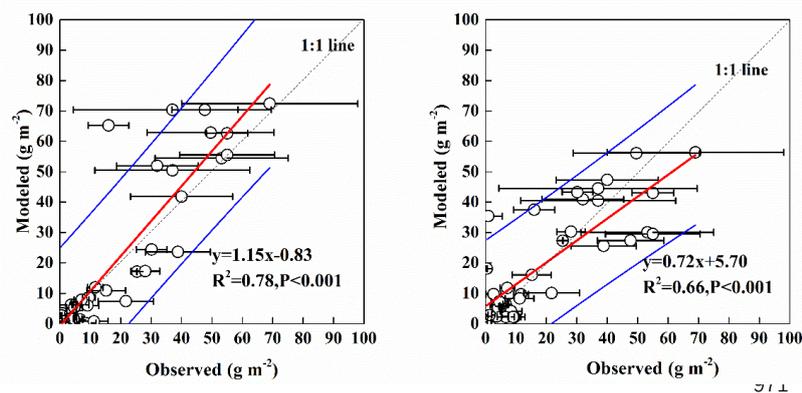
Continent/Wetland type	CH ₄ Flux (g m ⁻² yr ⁻¹)		CH ₄ MOD _{wetland}		CH ₄ Emissions (Tg)		Area (10 ⁴ × km ²)
	CH ₄ MOD _{wetland}	TEM	CH ₄ MOD _{wetland}	TEM	CH ₄ MOD _{wetland}	TEM	
Asia	23.27 ± 0.67	25.78 ± 0.14	25.37 ± 0.73	28.81 ± 0.15	109.04	904	902
Africa	27.64 ± 1.55	33.92 ± 0.27	21.12 ± 1.18	25.91 ± 0.20	76.39	905	905
N. America	11.48 ± 0.47	14.10 ± 0.18	24.38 ± 1.00	29.95 ± 0.38	212.35	906	906
S. America	29.61 ± 0.52	39.91 ± 0.54	26.24 ± 0.46	35.36 ± 0.48	88.60	907	907
Europe	7.77 ± 0.11	16.04 ± 0.28	6.48 ± 0.09	13.38 ± 0.23	83.40	908	908
Oceania	12.08 ± 2.52	11.31 ± 0.57	1.72 ± 0.27	1.61 ± 0.06	14.23	908	908
Lake, river*	15.57	15.57	27.32	27.32	175.46	909	909
Marsh	28.64 ± 1.06	39.60 ± 0.28	37.47 ± 1.39	51.80 ± 0.37	131.61	910	910
Peatland	1.99 ± 0.09	9.87 ± 0.27	0.42 ± 0.02	2.06 ± 0.06	21.00	911	911
Swamp	31.58 ± 0.57	45.59 ± 0.91	17.37 ± 0.32	25.08 ± 0.50	55.34	912	912
Coastal wetland	9.44 ± 0.25	11.63 ± 0.14	2.85 ± 0.08	3.51 ± 0.04	30.32	913	913
Intermittent wetland	18.49 ± 2.08	18.29 ± 0.19	5.81 ± 0.65	5.75 ± 0.06	31.60	914	914
No-specific wetland	10.21 ± 0.68	13.63 ± 0.25	14.07 ± 0.93	18.79 ± 0.34	138.67	914	914
Global	18.03 ± 0.49	23.00 ± 0.15	105.31 ± 2.72	134.31 ± 0.84	584.00	915	915

* IPCC Tier 1 method was used to estimate the CH₄ emissions from lakes and rivers. The CH₄ emission factor was from the IPCC (1996).

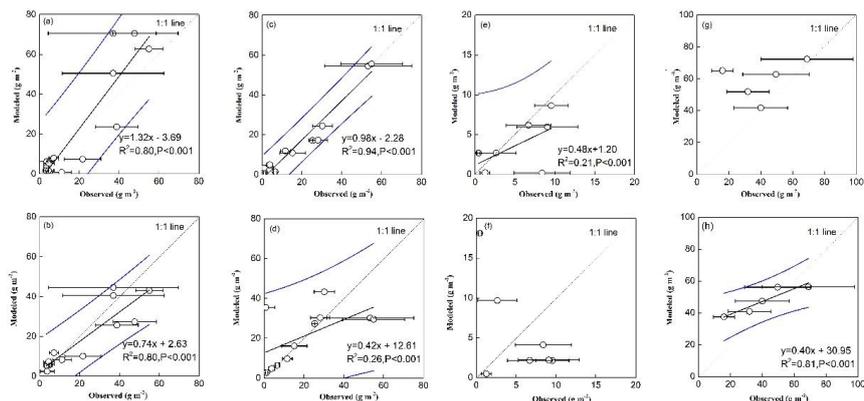


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932 **Figure 1: Wetland site distribution (Table 1) and global wetland maps of GLWD-3 (Lehner and Döll, 2004).**

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972 **Figure 2: Regression of simulated against observed total amount of seasonal CH₄ emission from global**
973 **wetland sites by CH4MOD_{wetland} (a) and the TEM (b). The horizontal bars are the standard errors from the**
974 **sampling replicates at the wetland site. The blue line is the prediction correspondence. The dashed line is the**
975 **1:1 line.**

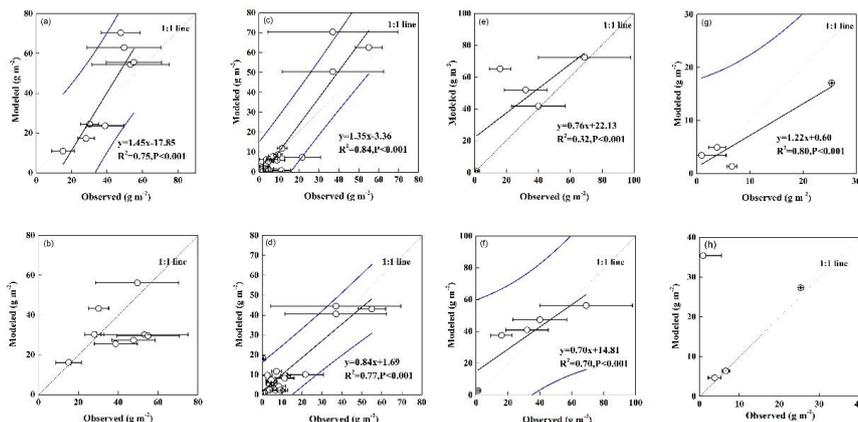


976 **Figure 3: Regression of simulated against observed total amount of seasonal CH₄ emissions from North**
977 **American wetland sites by CH4MOD_{wetland} (a) and the TEM (b), from Asian wetland sites by CH4MOD_{wetland}**
978 **(c) and the TEM (d), from European wetland sites by CH4MOD_{wetland} (e) and the TEM (f), and from South**
979 **American and African wetland sites by CH4MOD_{wetland} (g) and the TEM (h). The horizontal bars are the**
980 **standard errors from the sampling replicates at the wetland site. The blue line is the prediction**
981 **correspondence. The dashed line is the 1:1 line.**

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Figure 4: Regressions of simulated against observed total amount of seasonal CH₄ emissions from marsh sites by CH4MOD_{wetland} (a) and the TEM (b), from peatland sites by CH4MOD_{wetland} (c) and the TEM (d), from swamp sites by CH4MOD_{wetland} (e) and the TEM (f), and from coastal wetland sites by CH4MOD_{wetland} (g) and the TEM (h). The horizontal bars are the standard errors from the sampling replicates at the wetland site. The blue line is the prediction correspondence. The dashed line is the 1:1 line.

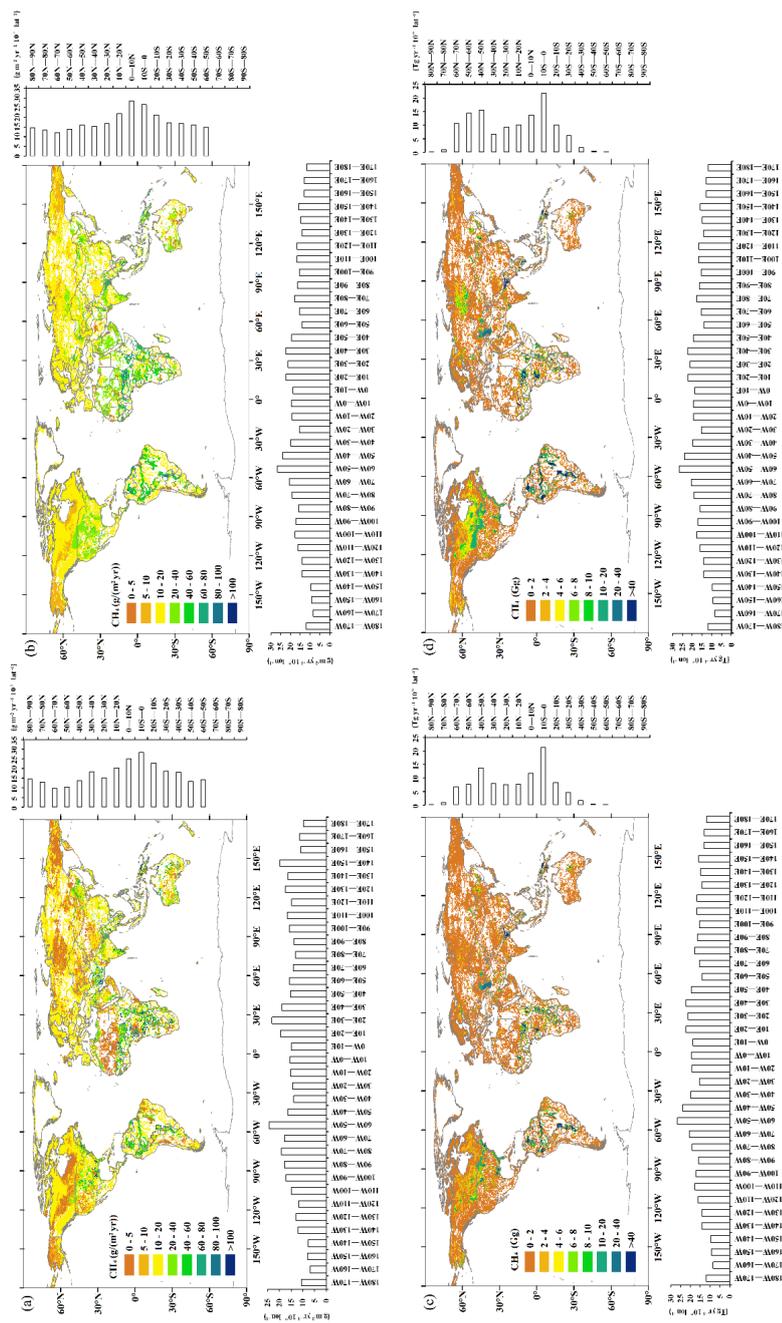


Figure 5: Spatial pattern of annual mean CH_4 fluxes for 2000–2010, with latitudinal and longitudinal distributions of annual mean CH_4 fluxes by $\text{CH}_4\text{MOD}_{\text{wetland}}$ (a) and the TEM (b).
 Spatial pattern of annual mean CH_4 emissions for 2000–2010, with latitudinal and longitudinal distributions of annual mean CH_4 emissions by $\text{CH}_4\text{MOD}_{\text{wetland}}$ (c) and the TEM (d).
 The CH_4 fluxes and emissions are aggregated in steps of 10° .