

1     **Evaluation of CH<sub>4</sub>MOD<sub>wetland</sub> and TEM models used to**  
2     **estimate global CH<sub>4</sub> emissions from natural wetlands**

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25 **Abstract**

26 Wetlands are the largest and most uncertain natural sources of atmospheric methane (CH<sub>4</sub>). Several  
27 process-based models have been developed to quantify the magnitude and estimate spatial and temporal  
28 variations in CH<sub>4</sub> emissions from global wetlands. Reliable models are required to estimate global  
29 wetland CH<sub>4</sub> emissions. This study aimed to test two process-based models, CH4MOD<sub>wetland</sub> and TEM,  
30 against the CH<sub>4</sub> flux measurements of marsh, swamp, peatland and coastal wetland sites across the world;  
31 specifically, model accuracy and generality were evaluated for different wetland types and in different  
32 continents, and then the global CH<sub>4</sub> emissions from 2000 to 2010 were estimated. Both models showed  
33 similar high correlations with the observed seasonal/annual total CH<sub>4</sub> emissions, and the regression of  
34 the observed versus computed total seasonal/annual CH<sub>4</sub> emissions resulted in R<sup>2</sup> values of 0.81 and 0.68  
35 for CH4MOD<sub>wetland</sub> and the TEM, respectively. The CH4MOD<sub>wetland</sub> produced accurate predictions for  
36 marshes, peatlands, swamps and coastal wetlands, with model efficiency (EF) values of 0.22, 0.52, 0.13  
37 and 0.72, respectively. The TEM produced good predictions for peatlands and swamps, with EF values  
38 of 0.69 and 0.74, respectively, but it could not accurately simulate marshes and coastal wetlands (EF<0).  
39 There was a good correlation between the simulated CH<sub>4</sub> fluxes and the observed values on most  
40 continents. However, CH4MOD<sub>wetland</sub> showed no correlation with the observed values in South America  
41 and Africa. The TEM showed no correlation with the observations in Europe. The global CH<sub>4</sub> emissions  
42 for the period 2000–2010 were estimated to be 105.31±2.72 Tg yr<sup>-1</sup> by CH4MOD<sub>wetland</sub> and 134.31±0.84  
43 Tg yr<sup>-1</sup> by the TEM. Both models simulated a similar spatial distribution of CH<sub>4</sub> emissions globally and  
44 on different continents. Marshes contribute 36–39% of global CH<sub>4</sub> emissions. Lakes/rivers and swamps  
45 are the second and third greatest contributors, respectively. Other wetland types account for only  
46 approximately 20% of global emissions. Based on the model applicability, if we use the more accurate  
47 model, i.e., the one that performs best as evidenced by a higher model efficiency and a lower model bias,  
48 to estimate each continent/wetland type, we obtain a new assessment of 116.99–124.74 Tg yr<sup>-1</sup> for the  
49 global CH<sub>4</sub> emissions for the period 2000–2010. Our results imply that performance at a global scale  
50 may conceal model uncertainty. Efforts should be made to improve model accuracy for different wetland  
51 types and regions, particularly hotspot regions, to reduce the uncertainty in global assessments.

## 52 **1 Introduction**

53 Atmospheric methane (CH<sub>4</sub>) is the second most prevalent human-induced greenhouse gas (GHG)  
54 after carbon dioxide (CO<sub>2</sub>). Its radiative forcing effect is 28 times greater than that of CO<sub>2</sub> on a 100-year  
55 horizon (Myhre et al., 2013). The radiative forcing attributed to CH<sub>4</sub> has been re-evaluated by the  
56 Intergovernmental Panel on Climate Change (IPCC)'s 5th Assessment Report (AR5) and was reported  
57 to be almost twice as high as the value reported in the 4th Assessment Report (AR4), with values of 0.97  
58 W m<sup>-2</sup> versus 0.48 W m<sup>-2</sup>, respectively (Myhre et al., 2013). This estimate considers that the emission of  
59 CH<sub>4</sub> leads to an increase in ozone production, stratospheric water vapor and CO<sub>2</sub>, which can affect its  
60 own lifetime (Boucher et al., 2009; Myhre et al., 2013; Shindell et al., 2012).

61 The growth rate of the atmospheric CH<sub>4</sub> concentration has varied in different historical periods.  
62 There was an exponential increase from preindustrial times to the 1980s. The growth rate decreased after  
63 the 1980s and was close to zero from 1999 to 2006; then, the growth rate resumed strong growth in the  
64 period of 2007-2017 (Dlugokencky et al., 2009, 2016; Nisbet et al., 2019). However, the causes that drive  
65 the variations in growth rate remain unclear due to the uncertainties in estimating CH<sub>4</sub> emissions and  
66 sinks (Ghosh et al., 2015; Saunio et al., 2016; Nisbet et al., 2019; Dalsøren et al., 2016).

67 Integrated at the global scale, wetlands are the largest and most uncertain source of CH<sub>4</sub> emitted to  
68 the atmosphere (Kirschke et al., 2013; Saunio et al., 2016). These emissions represent approximately  
69 30% of the total CH<sub>4</sub> input (Saunio et al., 2016). Bottom-up and top-down approaches are popular  
70 methods for estimating global CH<sub>4</sub> emissions from natural wetlands. Top-down approaches are based on  
71 inverse models (e.g., Bousquet et al., 2006; Fraser et al., 2013; Meirink et al., 2008; Tsuruta et al., 2017;  
72 Bruhwiler et al., 2014), which determine 'optimal' surface fluxes that best fit atmospheric CH<sub>4</sub>  
73 observations given an atmospheric transport model including chemistry, prior estimates of fluxes, and  
74 their uncertainties (Kirschke et al., 2013). Bottom-up approaches use process-based models that describe  
75 the relationship between the environmental factors and the processes of CH<sub>4</sub> production, oxidation and  
76 emission using mathematical equations (e.g., Li et al., 2010; Zhu et al., 2013; Zhang et al., 2002; Zhu et  
77 al., 2014; Walter and Heimann, 2000; Tian et al., 2015; Riley et al., 2011; Meng et al., 2012; Zhuang et  
78 al., 2006).

79 Recent studies related to the bottom-up approach have used an ensemble of process-based models  
80 driven by the same climate forcing to estimate the global CH<sub>4</sub> emissions from natural wetlands. For

81 example, the Wetland and Wetland CH<sub>4</sub> Intercomparison of Models Project (WETCHIMP) used ten land  
82 surface models and estimated global CH<sub>4</sub> emissions of  $190 \pm 76$  Tg CH<sub>4</sub> yr<sup>-1</sup> for the 1993–2004 period  
83 (Melton et al., 2013). In the following year, Kirschke et al. (2013) assessed a large emission range of  
84 142–287 Tg CH<sub>4</sub> yr<sup>-1</sup> from 1980 to 2010. Saunio et al. (2016) and Poulter et al. (2017) estimated global  
85 emissions of 153–227 Tg CH<sub>4</sub> yr<sup>-1</sup> for the decade 2003–2012 and  $184 \pm 22$  Tg CH<sub>4</sub> yr<sup>-1</sup> for the decade  
86 2000–2012 using ensemble process-based models (Poulter et al., 2017). Saunio et al. (2016) suggested  
87 that approximately 70% of the uncertainty was due to model structures and parameters.

88 Natural wetland ecosystems are greatly heterogeneous on a global scale. Wetlands vary widely by  
89 continent with respect to area and type (Kingsford et al., 2016; Keddy, 2010). Some wetland types have  
90 higher emissions, while some emit less CH<sub>4</sub>; this difference is because the processes of controls on CH<sub>4</sub>  
91 cycling differ among wetland types (Bridgman et al., 2013). For example, sedge-dominated marsh or fen  
92 often emit higher CH<sub>4</sub> fluxes because sedges can increase methanogenic substrates as part of their plant  
93 productivity and promote CH<sub>4</sub> transportation through their soft aerenchyma and lacunae tissues  
94 (McEwing et al., 2015; Jitka et al., 2017; Bhullar et al., 2013; Joabsson and Christensen, 2001; Kwon et  
95 al., 2017; King et al., 2002; Chanton, 2005). Bog soils with anaerobic incubations emit little CH<sub>4</sub> due to  
96 the particularly high CO<sub>2</sub>:CH<sub>4</sub> ratios of the end products of anaerobic carbon (Bridgman et al., 1998;  
97 Galand et al., 2010; Keller and Bridgman, 2007). Coastal wetlands with high salinity usually emit less  
98 CH<sub>4</sub> than other wetlands because the sulfate in seawater inhibits CH<sub>4</sub> production (Bartlett et al., 1985;  
99 Delaune et al., 1983; Li et al., 2016; Poffenbarger et al., 2011).

100 Model evaluation is a core part of model development and testing (Bennett et al., 2013). Based on  
101 the model evaluation, the modeler must be confident that the model will fulfill its purpose (Bennett et al.,  
102 2013; Rykiel, 1996). If applying process-based models for global-scale CH<sub>4</sub> estimations, it is necessary  
103 to evaluate its performance in different wetland types and regions. This process is also helpful for  
104 confirming the source of uncertainties and improving the model. However, previous studies have always  
105 focused on global assessments and have overlooked model performance in different wetland types or  
106 regions, which may have induced high uncertainties (Poulter et al., 2017; Saunio et al., 2016; Kirschke  
107 et al., 2013; Melton et al., 2013). CH<sub>4</sub>MOD<sub>wetland</sub> (Li et al., 2010) and the Terrestrial Ecosystem Model  
108 (TEM) (Zhuang et al., 2004; Melillo et al., 1993; Zhuang et al., 2007; Zhuang et al., 2013) are two  
109 established process-based models that can be used to simulate regional and global wetland CH<sub>4</sub> emissions.

110 Both models have been validated at specific sites (Zhu et al., 2013; Li et al., 2010; Li et al., 2017).  
111 However, we do not have information on the accuracy and applicability of the models for different  
112 wetland types and on different continents. The objectives of this study were to comprehensively evaluate  
113 the model performances of CH<sub>4</sub>MOD<sub>wetland</sub> and the TEM for different wetland types and on different  
114 continents and then to use the models to estimate global CH<sub>4</sub> emissions from natural wetlands.

## 115 **2 Methods and Materials**

116 The performance evaluation should clearly depend on the model objectives (Bennett et al., 2013).  
117 The models considered in this study aim to estimate the annual emissions from global wetlands.  
118 Therefore, the accuracy and applicability of the model in simulating annual CH<sub>4</sub> emissions for different  
119 wetland types and continents are very important in a performance evaluation. Several process-based  
120 models have been developed in recent decades (Xu et al., 2016). Some models are simple semiempirical  
121 models that focus on the biochemical processes of CH<sub>4</sub> production, oxidation and emission, e.g., Walter's  
122 model (Walter et al., 1996; Walter and Heimann, 2000), CASA (Potter, 1997) and CH<sub>4</sub>MOD<sub>wetland</sub> (Li et  
123 al., 2010). This kind of model requires simple inputs and parameters and is easily extrapolated to a  
124 regional scale. Other models are based on more complex land ecosystem models coupled to the CH<sub>4</sub>  
125 processes module, such as CLM4Me, ORCHIDEE, SDGVM and TEM. These models describe complex  
126 ecosystem processes require more inputs and parameters. In this study, we chose CH<sub>4</sub>MOD<sub>wetland</sub> and the  
127 TEM to compare the model performance of a simple easy-to-run model and a sophisticated land  
128 ecosystem model. Moreover, both models have been validated at the site scale, but no comprehensive  
129 accuracy analysis in different continents or for various wetland types has been done before. We collected  
130 CH<sub>4</sub> flux measurements from 43 wetlands spanning the main wetland types in the world from peer-  
131 reviewed literature (Table 1). A set of statistical methods was used to comprehensively evaluate the  
132 performance of CH<sub>4</sub>MOD<sub>wetland</sub> and the TEM in different wetland types and on different continents.  
133 Finally, we extrapolated both models to estimate the global CH<sub>4</sub> emissions from 2000 to 2010.

## 134 **2.1 Model overview**

### 135 **2.1.1 CH4MOD<sub>wetland</sub>**

136 The CH4MOD<sub>wetland</sub> model is a process-based biogeophysical model used to simulate the processes  
137 of CH<sub>4</sub> production, oxidation and emission from natural wetlands (Li et al., 2010). The model was  
138 established based on CH4MOD, which is used to predict CH<sub>4</sub> emissions from rice paddies (Huang et al.,  
139 1998; Huang et al., 1997). In CH4MOD<sub>wetland</sub>, we focused on the differences in the supply of  
140 methanogenic substrates between natural wetlands and rice paddies. Methanogenic substrates are derived  
141 from root exudates, the decomposition of plant litter and soil organic matter. The methane production  
142 rates were determined based on the methanogenic substrates and the influence of environmental factors,  
143 including soil temperature, soil texture and soil redox potential. Additionally, we incorporated the  
144 influence of salinity on CH<sub>4</sub> production to improve the model performance for coastal wetlands (Li et al.,  
145 2016). Inputs to the CH4MOD<sub>wetland</sub> model include the daily air/soil temperature, water table depth,  
146 annual aboveground net primary productivity (ANPP), soil sand fraction, soil organic matter, bulk  
147 density and soil salinity. The outputs are the daily and annual CH<sub>4</sub> production and emissions. We used  
148 the TOPMODEL hydrological model to simulate the water table depth as the inputs of CH4MOD<sub>wetland</sub>  
149 (Bohn et al., 2007; Li et al., 2015; Li et al., 2019; Zhu et al., 2013; Beven and Kirkby, 1979).

150 The main parameters that must be calibrated in CH4MOD<sub>wetland</sub> include the vegetation index (*VI*),  
151 which was used to quantify the different capacities for producing root exudates of the various plant  
152 species, the fraction of plant-mediated transport available (*T<sub>veg</sub>*), the fraction of CH<sub>4</sub> oxidized during  
153 plant-mediated transport (*P<sub>ox</sub>*), the proportion of belowground net primary productivity (*BNPP*) to the  
154 total net primary productivity (*NPP*) (*f<sub>r</sub>*), the fraction of nonstructural component in plant litter (*F<sub>N</sub>*)  
155 (Table S1) and the empirical constant of the influence of salinity. The model parametrization and main  
156 parameters are described in Supplementary Material S1.

### 157 **2.1.2 TEM**

158 The TEM is another process-based biogeochemical model that couples carbon, nitrogen, water, and  
159 heat processes in terrestrial ecosystems to simulate ecosystem carbon and nitrogen dynamics (Melillo et  
160 al., 1993; Zhuang et al., 2007; Zhuang et al., 2013). The methane dynamics module was first coupled

161 within the TEM by Zhuang et al. (2004) to explicitly simulate the process of methane production  
162 (methanogenesis), oxidation (methanotrophy) and transport between the soil and the atmosphere.  
163 Methane production is assumed to occur only in saturated zones and is regulated by organic substrate,  
164 soil thermal conditions, soil pH, and soil redox potentials; methane oxidation, which occurs in the  
165 unsaturated zone, depends on the soil methane and oxygen concentrations, temperature, moisture and  
166 redox potential. Methane transport is described by three pathways in the TEM: (1) diffusion through the  
167 soil profile, (2) plant-aided transport, and (3) ebullition. The TEM has also been coupled with  
168 TOPMODEL (Zhu et al., 2013). The model calibration of the TEM is well documented in Supplementary  
169 Material S2 and Table S2.

## 170 **2.2 Site information and data sources**

### 171 **2.2.1 Site information**

172 We collected 43 wetland sites across the world (Table 1). The wetland sites included 6 marsh sites,  
173 25 peatland sites, 8 swamp sites and 4 coastal wetland sites. Among the wetland sites, 7 sites are  
174 distributed in Europe (EU), 11 sites are distributed in Asia (AS), 2 sites are distributed in Africa (AF), 4  
175 sites are distributed in South America (SA) and 19 sites are distributed in North America (NA). The  
176 observations were from the late 1980s to the 2010s. The observation periods covered either a growing  
177 season or a whole year (Table 1). We calculated the total amount of CH<sub>4</sub> emissions during the growing  
178 season or the whole year as the observed seasonal/annual CH<sub>4</sub> emissions. For most of the wetland sites,  
179 the total amount of seasonal/annual CH<sub>4</sub> emissions during the observation period was calculated by  
180 summing the daily observations. Gaps in the CH<sub>4</sub> emission measurements were filled by linear  
181 interpolation between two adjacent days of observations. For a few wetland sites, the observed  
182 seasonal/annual CH<sub>4</sub> emissions were directly obtained from the literature. More details about the location,  
183 vegetation and observation periods are described in Table 1.

### 184 **2.2.2 Wetland map**

185 The global wetland distributions of different wetland types were based on the Global Lakes and  
186 Wetlands Database (GLWD-3) (<http://www.wwfus.org/science/data.cfm>) (Lehner and Döll, 2004) (Fig.

187 1). According to GLWD-3, the wetland types include 1. lakes, 2. reservoirs, 3. rivers (we combined lakes,  
188 reservoirs and rivers as a single wetland type, hereafter referred to as lakes/rivers), 4. freshwater marsh  
189 and floodplain (hereafter referred to as marsh), 5. swamp forest and flooded forest (hereafter referred to  
190 as swamp), 6. coastal wetland, 7. saline wetland (we combined coastal wetland and saline wetland as a  
191 single wetland type, hereafter referred to as coastal wetland), 8. bog, fen and mire (hereafter referred to  
192 as peatland), 9. intermittent wetland and 10. no-specific wetland. All of the observed sites (Table 1) are  
193 distributed on the wetland map (Fig. 1).

194 The global wetland area (excluding rivers) was estimated by the “Global Review of Wetland  
195 Resources and Priorities for Wetland Inventory (GRoWI)” as 530-570 M ha (Spiers, 1999). We used an  
196 average value, as the wetland area excluded rivers in this study. The global wetland area of rivers was  
197 based on GLWD-3. Therefore, we assumed that the global wetland area was 584 M ha, which represented  
198 the wetland area for the period from 2000 to 2010. The cartography-based GLWD-3 data provide a global  
199 distribution of natural wetlands at a 30-second resolution. Then, we aggregated the merged map up to  
200  $0.5^{\circ} \times 0.5^{\circ}$  (latitude  $\times$  longitude) grids. The wetland area (excluding rivers) in each pixel was adjusted  
201 by the ratio of the global wetland area estimated by GRoWI and by GLWD-3.

### 202 **2.2.3 Driver data**

203 The input climate data for the models include the daily air temperature, precipitation, cloudiness  
204 and vapor pressure. The historical daily climate data were developed from the latest monthly data sets of  
205 the Climatic Research Unit (CRU TS 3.10) of the University of East Anglia in the United Kingdom  
206 (Harris et al., 2014).

207 The soil properties needed by the CH4MOD<sub>wetland</sub> model include soil texture (percentage of sand in  
208 the soil), bulk density, soil organic carbon content, soil temperature and soil moisture. The additional  
209 information needed by the TEM includes the percentage of silt and clay in the soil, soil pH and site  
210 elevation. The soil texture data were derived from the soil map of the Food and Agriculture Organization  
211 (FAO) (FAO, 2012). The soil organic carbon content and the reference bulk density of wetland soils were  
212 retrieved from the Harmonized World Soil Database (HWSD) (FAO, 2008) by masking the HWSD with  
213 the Global Lakes and Wetlands Database (GLWD) (Lehner and Döll, 2004). The daily soil temperature  
214 data were estimated by the TEM from spatially interpolated climate data. The daily soil moisture driving



215 CH4MOD<sub>wetland</sub> coupled with TOPMODEL was developed from the monthly dataset  
216 ([http://www.cpc.ncep.noaa.gov/soilmst/leaky\\_glb.htm](http://www.cpc.ncep.noaa.gov/soilmst/leaky_glb.htm)) by temporal linear interpolation (Fan and van  
217 den Dool, 2004). The soil pH was also derived from the global soil property dataset of the International  
218 Geosphere-Biosphere Programme (IGBP) (Carter and Scholes, 2000).

219 The vegetation map of the IGBP was referenced to specify the vegetation parameters for  
220 CH4MOD<sub>wetland</sub> (Table S1) and the TEM. The map was derived from the IGBP Data and Information  
221 System (DIS) DISCover Database (Belward et al., 1999; Loveland et al., 2000). The 1 km × 1 km  
222 DISCover dataset was reclassified into the TEM vegetation classification scheme and then aggregated  
223 into 0.5° × 0.5° grids. The annual ANPP used to drive CH4MOD<sub>wetland</sub> was from the output of the TEM.

224 For CH4MOD<sub>wetland</sub>, a high-resolution topographic wetness index dataset (Marthews et al., 2015)  
225 was used to calculate the changes in the water table. Global salinity data were obtained from the World  
226 Ocean Atlas 2009 (Antonov et al., 2010). We also used 1 km × 1 km global elevation data derived from  
227 the Shuttle Radar Topography Mission (SRTM) (Farr et al., 2007). The above data were resampled to  
228 0.5° × 0.5° grids to match the resolution of the other input data.

### 229 **2.3 Model evaluation**

230 We compared the observed seasonal/annual CH<sub>4</sub> emissions from the wetland sites (Table 1) and the  
231 simulated CH<sub>4</sub> emissions at the 0.5°×0.5° grid scale for the same period (described in Sect. 2.4). The  
232 statistics include the determination coefficient (R<sup>2</sup>), the root mean square error (RMSE), the mean  
233 deviation (RMD), the model efficiency (EF) and the coefficient of determination (CD) were used to  
234 evaluate model performance on a global scale, a continental scale and for each wetland type. Because of  
235 the limited number of sites in Africa and South America, we combined the two continents together.

236 Two simulations with the same RMSE values may not be considered equivalent because the  
237 distribution of the error among the sources may not be the same (Allen and Raktoc, 1981). We further  
238 analyzed the source of the model errors by decomposing it into three components: the mean bias from  
239 the modeling procedure (U<sub>M</sub>), the errors due to regression (U<sub>R</sub>) and the errors due to random disturbances  
240 (U<sub>E</sub>) (Allen and Raktoc, 1981). The detailed description and the equations used to calculate these  
241 statistics are described in Supplementary Material S3.

## 242 **2.4 Model extrapolation**

243 CH4MOD<sub>wetland</sub> and the TEM were used to simulate the CH<sub>4</sub> emissions from global wetlands at a  
244 spatial resolution of 0.5°×0.5°. We established spatially explicit data for climate, soils, vegetation, land  
245 use and other environmental inputs at a 0.5°×0.5° spatial resolution to facilitate the models at the global  
246 scale. Both process-based models were conducted for the period of 1980–2010 in each pixel to simulate  
247 the temporal spatial variations in CH<sub>4</sub> fluxes. In this study, we focused only on the total CH<sub>4</sub> emissions  
248 for the period 2000–2010 because we assumed that the wetland map represented the distribution of  
249 natural wetlands during this period. The total CH<sub>4</sub> emissions from the natural wetlands, excluding the  
250 lakes/rivers in each pixel, were calculated as the product of the CH<sub>4</sub> fluxes and the gridded wetland area.  
251 To make an overall global/continental CH<sub>4</sub> emissions assessment, we evaluated the CH<sub>4</sub> emissions from  
252 lakes/rivers using the IPCC Tier 1 method based on the CH<sub>4</sub> emissions factor (IPCC, 1996) and the area  
253 of lakes/rivers in each pixel.

254 We aggregated the gridded values and obtained the annual mean CH<sub>4</sub> emissions from each wetland  
255 type and each continent by CH4MOD<sub>wetland</sub> combined with the IPCC Tier1 method (hereafter referred to  
256 as Method A) and the TEM combined with the IPCC Tier1 method (hereafter referred to as Method B).  
257 In addition, the two global assessments Method A and Method B, we made two other assessments of  
258 global CH<sub>4</sub> emissions by choosing the more accurate model (Method C and Method D). Based on the  
259 model performance evaluation (Sect. 2.3), we found a more accurate model for each wetland type and  
260 each continent. In the Method C approach, we chose the CH<sub>4</sub> emissions from each continent simulated  
261 by the more accurate model. In Oceania, we used the average simulated result by CH4MOD<sub>wetland</sub> and  
262 the TEM because there was no wetland site on this continent (Table 1). We summed the CH<sub>4</sub> emissions  
263 from all continents and made an assessment of the global CH<sub>4</sub> emissions. In the Method D approach, we  
264 chose the CH<sub>4</sub> emissions from marsh, peatland, swamp and coastal wetlands simulated by the more  
265 accurate model. The CH<sub>4</sub> emissions from intermittent wetlands and nonspecific wetlands were used as  
266 the average result by CH4MOD<sub>wetland</sub> and the TEM. The CH<sub>4</sub> emissions from lakes/rivers were based on  
267 the IPCC Tier 1 method. We summed the CH<sub>4</sub> emissions from all wetland types and assessed the global  
268 CH<sub>4</sub> emissions.

## 269 3. Results

### 270 3.1 Model evaluation

#### 271 3.1.1 Model evaluation for global wetland sites

272 Fig. 2 shows the correlation of the modeled versus observed total amount of seasonal/annual CH<sub>4</sub>  
273 emissions by CH4MOD<sub>wetland</sub> (Fig. 2a) and the TEM (Fig. 2b). The regression of the observed versus  
274 computed total seasonal/annual CH<sub>4</sub> emissions by CH4MOD<sub>wetland</sub> (Fig. 2a) resulted in an R<sup>2</sup> of 0.81,  
275 with a slope of 1.17 and an intercept of -1.93 g m<sup>-2</sup> (n=58, p<0.001). The regression of the observed  
276 versus computed total seasonal/annual CH<sub>4</sub> emissions by the TEM (Fig. 2b) resulted in an R<sup>2</sup> of 0.68,  
277 with a slope of 0.74 and an intercept of 4.77 g m<sup>-2</sup> (n=58, p<0.001). These results indicated that the  
278 variations in the CH<sub>4</sub> emissions between sites and in different years could be delineated by both process-  
279 based models.

280 The statistics of the model performance of seasonal/annual CH<sub>4</sub> emissions (Table 2) indicated that  
281 both process-based models had the capability to simulate seasonal/annual CH<sub>4</sub> emissions from natural  
282 wetlands on a global scale (EF=0.65 for CH4MOD<sub>wetland</sub> and EF=0.68 for the TEM). However, a  
283 discrepancy still existed between the simulated and observed seasonal/annual CH<sub>4</sub> emissions  
284 (RMSE=67.00% for CH4MOD<sub>wetland</sub> and RMSE=63.58% for the TEM). For CH4MOD<sub>wetland</sub>, the source  
285 of the errors was mainly from the regression error and random error, while for the TEM, the errors were  
286 mainly due to random disturbances (Table 2). Both models slightly overestimated the seasonal/annual  
287 CH<sub>4</sub> emissions on a global scale, with RMD values of ~4% (Table 2).

#### 288 3.1.2 Model evaluation for different continents

289 We further analyzed the model predictions by CH4MOD<sub>wetland</sub> and the TEM among different  
290 continents (Fig. 3, Table 2). There was a good correlation between the simulated seasonal/annual CH<sub>4</sub>  
291 emissions and the observed values on most of the continents by the two models. The R<sup>2</sup> varied between  
292 0.35 (Fig. 3e) and 0.94 (Fig. 3c) for CH4MOD<sub>wetland</sub> and between 0.26 (Fig. 3d) and 0.80 (Fig. 3h) for  
293 the TEM. The CH4MOD<sub>wetland</sub> model yielded more accurate predictions in Asia and North America, with  
294 EFs of 0.93 and 0.57, respectively (Fig. 3b and 3a, Table 2), than in South America and Africa (EF <0 in  
295 Table 2) (Fig. 3g). The TEM yielded more accurate predictions in North America and South

296 America/Africa than  $CH_4MOD_{wetland}$ , with EF values of 0.76 and 0.53, but performed poorly in Europe  
297 (EF <0 in Table 2).  $CH_4MOD_{wetland}$  underestimated the observed emissions (RMD = -12.64%) in Asia  
298 and Europe (RMD = -29.91%) (Table 2). The TEM overestimated the  $CH_4$  emissions in South  
299 America/Africa (RMD=15.31%) and slightly underestimated the  $CH_4$  emissions in North America  
300 (RMD=-2.86%) (Table 1). Random error was the main contributor to the model errors in Asia and Europe  
301 in  $CH_4MOD_{wetland}$  and in Asia, North America, and South America/Africa in the TEM (Table 2). However,  
302 the regression error contributed most to the model errors in North America in  $CH_4MOD_{wetland}$  (Table 2).

### 303 **3.1.3 Model evaluation for different wetland types**

304 Fig. 4 shows the regressions of the simulated values against the observed total amount of  
305 seasonal/annual  $CH_4$  emissions from the different wetland types. Regression analysis indicated that both  
306 models showed good performance in modeling seasonal/annual  $CH_4$  emissions from the peatland sites  
307 (Fig. 4c and d). The TEM showed a better model efficiency and a lower RMSE and RMD than the  
308  $CH_4MOD_{wetland}$  (Table 2) for peatland. For the other wetland types,  $CH_4MOD_{wetland}$  showed good  
309 performance in simulating the seasonal/annual  $CH_4$  emissions from coastal wetlands (EF = 0.72),  
310 followed by marshes (EF = 0.22) and swamps (EF = 0.13) (Table 2). The TEM showed poor performance  
311 for the marsh sites (EF = -0.42) and coastal wetlands (EF = -2.26) (Table 2); however, it showed good  
312 performance for the swamp sites (EF = 0.74). There was no significant correlation ( $p>0.05$ ) between the  
313 modeled and observed seasonal/annual  $CH_4$  emissions from the marsh sites (Fig. 4b) and coastal wetland  
314 sites (Fig. 4h).

315 The errors by  $CH_4MOD_{wetland}$  were mainly due to the regression error for marsh and peatland (Table  
316 2). For coastal wetlands, the model bias contributed 24%, the regression error contributed 30%, and the  
317 random error contributed 47% to the model errors (Table 2). The errors by the TEM were mainly due to  
318 the random error in peatland and swamps (Table 2).

## 319 **3.2 Global $CH_4$ emissions from natural wetlands**

### 320 **3.2.1 Spatial pattern of global $CH_4$ emissions**

321

322 The distribution of the simulated annual mean CH<sub>4</sub> fluxes and total CH<sub>4</sub> emissions for the period  
323 2000–2010 showed similar patterns in CH4MOD<sub>wetland</sub> and the TEM (Fig. 5). The simulated latitudinal  
324 contributions of CH<sub>4</sub> fluxes were consistent between the two models (Supplementary material S4, Fig.  
325 5a and 5b). Large emissions were found in South America, southern Africa and the border of Canada and  
326 the United States (Fig. 5c and 5d). The latitudinal sums of CH<sub>4</sub> emissions indicated that the strongest  
327 contribution came from the tropical zone (Fig. 5c and 5d). The latitudinal band of 10°S–0° contributed  
328 22.77 Tg yr<sup>-1</sup> and 23.23 Tg yr<sup>-1</sup> CH<sub>4</sub> in CH4MOD<sub>wetland</sub> and the TEM, which accounted for 22% and 18%  
329 of the global emissions, respectively. A secondary large peak was simulated in the 40–50°N latitudinal  
330 band, with a value of 14.64 Tg yr<sup>-1</sup> and 16.66 Tg yr<sup>-1</sup> CH<sub>4</sub> according to CH4MOD<sub>wetland</sub> and the TEM,  
331 respectively. Generally, both models simulated a common decline in CH<sub>4</sub> emissions from lower latitudes  
332 to higher latitudes (Fig. 5c and 5d). The largest peak in CH<sub>4</sub> emissions was modeled in the 60–50°W  
333 meridional band, with values of 11.63 Tg yr<sup>-1</sup> in CH4MOD<sub>wetland</sub> (Fig. 5c) and 13.83 Tg yr<sup>-1</sup> in the TEM  
334 (Fig. 5d). This peak corresponded to the longitudes of the Amazon in South America. Both models  
335 simulated secondary peaks in the 30–40°E meridional band (Fig. 5c and 5d), which corresponded to the  
336 longitudes of the Congo in Africa.

### 337 **3.2.2 CH<sub>4</sub> emissions from different continents and wetland types**

338 Table 3 provides an overview of the CH<sub>4</sub> emissions from different continents and wetland types  
339 simulated by CH4MOD<sub>wetland</sub> and the TEM. A comparison of simulated CH<sub>4</sub> fluxes from different  
340 continents by CH4MOD<sub>wetland</sub> and the TEM showed that the three highest fluxes were modeled in South  
341 America, Africa and Asia (Table 3). The TEM simulated higher CH<sub>4</sub> fluxes in Europe than in North  
342 America, but the CH4MOD<sub>wetland</sub> simulations showed the opposite. For Oceania, the two models  
343 simulated similar fluxes.

344 Both models simulated the same sequence of CH<sub>4</sub> fluxes: swamp, marsh, intermittent wetland, no-  
345 specific wetland, coastal wetland, and peatland (Table 3). The simulated annual mean CH<sub>4</sub> fluxes from  
346 intermittent wetlands were almost equivalent in both models. For other wetland types, the TEM simulated  
347 higher CH<sub>4</sub> fluxes than the CH4MOD<sub>wetland</sub> model (Table 3). Both models simulated peak emissions in  
348 summer and lower emissions in winter for all wetland types except swamps (Fig. S1). Since large area  
349 of swamps distributed in southern hemisphere (Fig. 1), higher and lower CH<sub>4</sub> emissions were simulated

350 during March to May and June to August, respectively (Fig. S1).

351 The global CH<sub>4</sub> emissions simulated by the TEM were 29 Tg yr<sup>-1</sup> higher than those simulated by  
352 CH4MOD<sub>wetland</sub> (Table 3). This difference depended on the differences in the CH<sub>4</sub> fluxes and on the  
353 wetland area. The simulated results showed that half of this difference was attributed to marshes. South  
354 America contributed 30% to this difference because the simulated CH<sub>4</sub> fluxes differed greatly between  
355 the TEM and CH4MOD<sub>wetland</sub> (Table 3).

356 The two models simulated similar spatial distributions of the CH<sub>4</sub> emissions among different  
357 wetland types and continents (Table 3). Marshes emit higher CH<sub>4</sub> fluxes and have the largest area. Thus,  
358 marshes were the greatest contributor to global CH<sub>4</sub> emissions and contributed 36%–39% to global CH<sub>4</sub>  
359 emissions (Table 3). Lakes/rivers and swamps were the second and third contributors, respectively (Table  
360 3). The CH<sub>4</sub> emissions from peatlands, coastal wetlands, intermittent wetlands and no-specific wetlands  
361 accounted for only approximately 20% of the global emissions (Table 3).

362 Although North America accounted for 36% of the global wetland area, it contributed only 22%–  
363 23% to global emissions (Table 3). In contrast, the wetland area in South America accounted for 15% of  
364 the global area and contributed 25%–26% to global CH<sub>4</sub> emissions. Asia and Africa also accounted for  
365 approximately 20% of global emissions. The lowest area and emissions were found in Oceania (Table  
366 3).

### 367 **3.2.3 Global CH<sub>4</sub> estimations**

368 The global CH<sub>4</sub> emissions for the period 2000-2010 were estimated to be 105.31 ± 2.72 Tg yr<sup>-1</sup> by  
369 Method A and 134.31 ± 0.84 Tg yr<sup>-1</sup> by Method B. Based on the evaluation of model performance (Table  
370 2), CH4MOD<sub>wetland</sub> yielded the most accurate predictions for Asia and Europe, and the TEM yielded the  
371 most accurate predictions for North America and South America/Africa. Using this combination, the  
372 global CH<sub>4</sub> emissions were estimated to be 124.74 ± 1.22 Tg by Method C. Similarly, in Method D,  
373 CH4MOD<sub>wetland</sub> was used for simulations in marshes and coastal wetlands, and the TEM was used for  
374 simulations in peatlands and swamps; as a result, the global wetland CH<sub>4</sub> emissions were estimated to be  
375 116.99 ± 2.23 Tg.

## 376 4. Discussion

### 377 4.1 Generality of CH<sub>4</sub>MOD<sub>wetland</sub> and the TEM

378 A lack of correspondence between the model output and observations could be partly due to the  
379 observed flux data, e.g., the inevitable gap-filling of missing data points to determine the seasonal/annual  
380 total emissions (Kramer et al., 2002). The results showed differences between the observed and simulated  
381 CH<sub>4</sub> emissions by both CH<sub>4</sub>MOD<sub>wetland</sub> and the TEM on a global scale (Fig. 2) and continent scale (Fig.  
382 3) and for different wetland types (Fig. 4). The reliability of the observed flux data was not under  
383 discussion in this study. We evaluated only the model accuracy and applicability across different wetland  
384 types and continents.

385 On a global scale, both models fulfilled the criteria of sufficient accuracy for the ability to predict  
386 CH<sub>4</sub> fluxes (Table 2). However, this fuzzy analysis may miss some real model performance. For the  
387 model applicability on different continents, CH<sub>4</sub>MOD<sub>wetland</sub> performed best in Asia, followed by North  
388 America and Europe. It performed poorly in South America/Africa, where swamps are more common  
389 (Table 2). The TEM performed best in North America, followed by South America/Africa and Asia. It  
390 performed poorly in Europe (Table 2). Each continent has different main wetland types; thus, the model  
391 applicability for different continents depended on its applicability for different types. CH<sub>4</sub>MOD<sub>wetland</sub> is  
392 suitable for marshes, peatlands and coastal wetlands, but it cannot be applied in swamps (Table 2). This  
393 limitation may be because in CH<sub>4</sub>MOD<sub>wetland</sub>, only a semiempirical logistic model is used to simulate  
394 plant growth (Li et al., 2010). This characteristic may induce large uncertainties in simulating the growth  
395 of forests in swamps (Table 1). However, the TEM uses the carbon nitrogen dynamics module (CNDM)  
396 to describe the effects of photosynthesis, respiration, decomposition and nutrient cycling on NPP (Melillo  
397 et al., 1993). Compared with CH<sub>4</sub>MOD<sub>wetland</sub>, the TEM performed well in simulating NPP in various  
398 vegetation types (Melillo et al., 1993). According to the model evaluation, the TEM was suitable for  
399 swamps and peatlands but had large uncertainties in marshes and coastal wetlands (Table 2). This pattern  
400 may be because the TEM focuses on two major wetland types: boreal tundra and forest wetland (Zhuang  
401 et al., 2004). The biochemical processes in the TEM model may be suitable for peatlands (tundra) and  
402 swamps (forest wetland) but not suitable for marshes. For coastal wetlands, the TEM did not consider  
403 the inhibition of salinity on CH<sub>4</sub> production (Poffenbarger et al., 2011; Bartlett et al., 1987) and greatly

404 overestimated the CH<sub>4</sub> fluxes (Table 2). CH4MOD<sub>wetland</sub> introduced the influence of salinity on CH<sub>4</sub>  
405 production and had good performance for coastal wetlands (Table 2).

#### 406 **4.2 Reducing uncertainties in global estimations**

407 The estimates of global wetland CH<sub>4</sub> emissions had large ranges in previous studies (Zhu et al.,  
408 2015). The estimates by process-based models ranged from 92 Tg yr<sup>-1</sup> (Cao et al., 1996) to 297 Tg yr<sup>-1</sup>  
409 (Gedney et al., 2004) during the period of 1980-2012. Recently, an ensemble of process-based models  
410 driven by the same climatic data has commonly been used to estimate global wetland CH<sub>4</sub> emissions  
411 (Melton et al., 2013; Kirschke et al., 2013; Poulter et al., 2017; Saunois et al., 2016). However, the  
412 uncertainties in the model mean estimation range from 12% (Poulter et al., 2017) to 40% (Melton et al.,  
413 2013). The uncertainty mainly comes from the wetland distribution and model structure and parameters  
414 (Saunois et al., 2016). Estimating accurate wetland extent and its seasonal and annual variations is a  
415 major challenge in present studies. The global estimations of wetland area ranged from 4.3 M ha to 12.9  
416 M ha during the period of 1990 to 2005 (Melton et al., 2013). The wetland extent of 9.2 M ha from the  
417 GLWD excluded water bodies, and this value was ~40% higher than the wetland area used in this study.  
418 That is, this difference was the main reason for the lower global estimations determined in this study than  
419 those reported in previous works (Zhu et al., 2015; Melton et al., 2013; Poulter et al., 2017; Saunois et  
420 al., 2016). Improving the accuracy of wetland extent and temporal variations is important in reducing  
421 uncertainties in global wetland CH<sub>4</sub> estimations.

422 In addition to wetland area, the model structure and parameters accounted for ~70% of the total  
423 uncertainties (Saunois et al., 2016). The results of the accuracy analysis showed that for CH4MOD<sub>wetland</sub>,  
424 regression bias accounted for 61% of the model errors in peatland, and mean bias accounted for 22% of  
425 the RMSE in swamp; for the TEM, mean bias and regression bias accounted for 29% and 42%,  
426 respectively, of the model errors in coastal wetland (Table 2). This result indicated that there were still  
427 uncertainties in the modeling procedure, e.g., in the model mechanism or in parameterization (Zhang et  
428 al., 2017; Allen and Raktoc, 1981). In the existing process-based models, which are not limited to  
429 CH4MOD<sub>wetland</sub> and the TEM, some important procedures should be focused on to reduce the bias due  
430 to the model mechanism. For example, the mechanism of the freeze-thaw cycle is important in process-  
431 based models (Wei and Wang, 2017) because of the large contribution of CH<sub>4</sub> released during the



432 nongrowing season in some frozen regions (Friborg et al., 1997; Huttunen et al., 2003; Mastepanov et  
433 al., 2008; Zona et al., 2016). In addition, quantifying CH<sub>4</sub> ebullition is important but difficult due to the  
434 uncertainty in estimates of CH<sub>4</sub> emissions from peatlands (Stanley et al., 2019). Moreover, although the  
435 importance of plants in CH<sub>4</sub> biogeochemical processes has been reported in many studies, better  
436 modeling and characterization of plant community structure is needed (Bridgham et al., 2013). Finally,  
437 most of the present process-based models do not have the ability to simulate CH<sub>4</sub> exchange from water  
438 bodies, such as lakes, rivers and reservoirs, although such water bodies contribute significantly to the  
439 global budget (Deemer et al., 2016). The use of the IPCC Tier method inevitably induces large  
440 uncertainties in the global estimates. The above mechanisms should be incorporated into existing  
441 process-based models to reduce the uncertainties in the current assessment.

442 The observational data related to processes of and controls on CH<sub>4</sub> production, consumption, and  
443 transport also limit the model calibration and validation. The flux data of 43 wetland sites used for model  
444 performance in this study are quite limited and do not represent all climatic, soil, hydrologic and  
445 vegetation conditions across global natural wetlands (Table 1). The observations in this study used both  
446 the chamber method and the eddy covariance method (Aubinet et al., 2012), which are widely used for  
447 CH<sub>4</sub> observations (Table 1). There are differences in measuring CH<sub>4</sub> fluxes between the two methods  
448 (Chaichana et al., 2018). The eddy covariance method may underestimate the fluxes (Twine et al., 2000;  
449 Sachs et al., 2010), while the chamber method may overestimate the fluxes (Werle and Kormann, 2001).  
450 These differences may introduce uncertainties to model calibration and validation. Furthermore, both  
451 process-based models were evaluated on an annual basis rather than on a daily scale. The validation of  
452 seasonal variation was not performed in this study, partly because we cannot obtain the daily step data  
453 for some of the sites. Fine temporal validation against more flux datasets, especially fluxes by eddy  
454 covariance experiments, and intermediate variables that control the CH<sub>4</sub> process are necessary in future  
455 studies (Wei and Wang, 2017).

## 456 **5. Conclusion**

457 Two process-based models, CH<sub>4</sub>MOD<sub>wetland</sub> and the TEM, were used to simulate annual CH<sub>4</sub>  
458 emissions from different wetland types and continents, and their performances were evaluated. Model  
459 validation showed that both models could simulate variations between different wetland sites and years.

460 The statistical analysis of model performance showed that CH4MOD<sub>wetland</sub> was capable of simulating  
461 CH<sub>4</sub> emissions from marshes, peatlands, swamps and coastal wetlands, while the TEM was capable of  
462 simulating CH<sub>4</sub> emissions from peatlands and swamps (model efficiency > 0). CH4MOD<sub>wetland</sub> performed  
463 well in Asia, Europe and North America, while the TEM performed well in North America, Asia, South  
464 America and Africa. The models were then used to estimate global wetland CH<sub>4</sub> emissions. The CH<sub>4</sub>  
465 simulations of both models had good agreement in terms of the latitudinal and meridional bands. The  
466 global CH<sub>4</sub> emissions for the period 2000–2010 were estimated to be 105.31± 2.72 Tg yr<sup>-1</sup> by  
467 CH4MOD<sub>wetland</sub> and 134.31 ± 0.84 Tg yr<sup>-1</sup> by the TEM. If we used a more accurate model to estimate  
468 each continent/wetland type based on the models' generality, the estimated global CH<sub>4</sub> emissions were  
469 116.99–124.74 Tg yr<sup>-1</sup> for the period 2000–2010. The uncertainty in global wetland CH<sub>4</sub> assessments by  
470 the process-based model approach comes from the inaccuracy of the wetland mapping area, the modeling  
471 procedure and the observational limitations. Future research on accurately mapping wetlands, improving  
472 model mechanisms and parametrization and using more observations to evaluate model performance  
473 would improve global estimations.

#### 474 **Code and data availability.**

475 The TEM and CH4MOD<sub>wetland</sub> model code and model data sets (input data and model results) are  
476 available on the website <https://zenodo.org/record/3594988#.XglabvkzY2w>.

#### 477 **Author Contribution**

478 T. Li and L. Yu pondered the rationale of the method. T. Li and Y. Lu developed and performed the model  
479 simulations. W. Sun, Q. Zhang, W. Zhang, G. Wang, Z. Qin, L. Yu, H. Li and R. Zhang made the data  
480 collection and processing. T. Li prepared the manuscript with contributions from all coauthors.  
481

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#### 485 **Competing Interests**

486 The authors declare no competing interests.

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**Table 1. Description of observation wetland sites**

ID	Wetland Name, Continent	Location	Wetland Type	Plant Species	Observation Period	Reference
1	Northeast Siberia, Russia, EU	72°22'N, 126°28'E	Peatland <sup>a</sup>	<i>Carex</i> spp., <i>Limprichtia revolvens</i> , <i>Meesia longiseta</i>	1999.5–1999.9 2003.7–2004.7 *	Wagner et al., 2003 Wille et al., 2008
2	Northeast Siberia, Russia, EU	71°30'N, 130°00'E	Peatland <sup>a</sup>	<i>Eriophorum</i> , <i>Carex</i> spp., <i>Sphagnum</i> spp., <i>Salix</i> spp.	1993.7–1993.8	Nakano et al., 2000
3	Northeast Siberia, Russia, EU	68°30'N, 161°24'E	Peatland <sup>a</sup>	<i>Larix</i> , <i>Alnus</i> spp., <i>Betula</i> spp., <i>Salix</i> spp.	1995.7–1995.8	Nakano et al., 2000
4	Northeast Siberia, Russia, EU	70°50'N, 147°29'E	Peatland <sup>a</sup>	<i>Betula nana</i> , <i>Salix pulchra</i> dwarf shrubs, sedge, <i>Sphagnum</i>	2008.7–2008.8 * 2009.6–2009.8 *	Parmentier et al., 2011
5	Zackenbergl, Greenland, EU	74°30'N, 21°00'W	Peatland	<i>Cassiope tetragona</i> , <i>Salix arctica</i>	1996.6–1996.8 1999.7–1999.8 2000.7–2000.8	Christensen et al., 2000; Joabsson and Christensen, 2001
6	Abisko, Sweden, EU	68°22'N, 19°03'E	Peatland <sup>a</sup>	<i>Eriophorum angustifolium</i> , <i>Carex</i> spp.	1974.6–1974.9 2008–2009 *	Svensson and Rosswall, 1984 Olefelt et al., 2012
7	Kaamanen, Finland, EU	69°08'N, 27°17'E	Peatland	Shrubs, <i>Carex</i> spp., moss, etc.	1998.4–1999.4 *	Aurela et al., 2002
8	Sanjiang Plain, China, AS &	47°35'N, 133°31'E	Marsh	<i>Carex lasiocarpa</i> , <i>Deyeuxia angustifolia</i>	2002.6–2005.11	Hao, 2006; Song et al., 2008
9	Ruoergai Plateau, China, AS	32°47'N, 102°32'E	Peatland	<i>Carex muliensis</i> , <i>Carex meyeriana</i>	2001.4–2001.10	Wang et al., 2002
10	Wuliangsu Lake, China, AS &	40°47'–41°03' N, 108°43'–108°57' E	Marsh	<i>Phragmites australis</i>	2003. 4–2003.10	Duan et al., 2005
11	Haibei alpine marsh, China, AS	37°29'N, 101°12'E	Marsh	<i>Carex allivescens</i>	2002.7–2002.9	Hirota et al., 2004
12	Zhalong Wetland, China, AS	46°52'N–47°32'N, 123°47'E–124°37'E	Marsh	<i>Phragmites australis</i>	2009.5–2009.10	Huang et al., 2011
13	Liao River delta, China, AS	40°40'–41°25'N, 121°35'–122°55'E	Coastal <sup>b</sup>	<i>Phragmites australis</i>	1997.4–1997.11	Huang et al., 2005
14	Chongming Island, China, AS &	31°15'N, 121°30'E	Coastal <sup>b</sup>	<i>Scirpus</i>	2004. 5–2004.12 2011.2–2011.12	Li et al., 2014
15	Guangzhou, China, AS	23°01'N, 113°46'E	Coastal <sup>c</sup>	<i>Aegiceras corniculatum</i> etc.	2005.3–2005.12 *	Kang et al., 2008
16	Haikou, China, AS	19°51' N, 110°24' E	Coastal <sup>c</sup>	<i>Bruguiera sexangula</i>	1996.1–1997.12 *	Ye et al., 2000
17	Sarawak, Malaysia, AS &	2°49'N, 111°51'E	Swamp	Flooded forest <sup>s</sup>	2002.8–2003.7	Melling et al., 2005
18	Kalimantan, Indonesia, AS	2°20'S, 113°55'E	Swamp	<i>Shorea balangerana</i>	1994.9–1995.9	Page et al., 1999; Jauhiainen et al., 2005
19	Congo River basin, Congo, AF	4°00'S–0°00', 14°00'–18°00'E	Swamp	Flooded forest <sup>s</sup>	1988 *	Tathy et al., 1992
20	Congo River basin, Congo, AF	0°00'–4°00'N, 14°00'–18°00'E	Swamp	Flooded forest <sup>s</sup>	1988 *	Tathy et al., 1992
21	Pantanal, Brazil, SA	19°30'S, 57°00'W	Marsh	<i>Paspalum repens</i>	1998.1–1998.12	Alvalá and Kirchoff, 2000; Melack et al., 2004

22	Lago Calado, Brazil, SA	3°15'S, 60°34'W	Swamp	Flooded forest <sup>§</sup>	1985 *	Crill et al., 1988
23	Central Brazilian Amazon, SA	5°00'S–0°00', 50°00'–70°00'W	Swamp	Flooded forest <sup>§</sup>	1985 *	Devol et al., 1988
24	Negro River basin, Brazil, SA	0°17'S, 63°34'W	Swamp	<i>Emergent sedge, shrub, palm</i>	2005.1–2006.1 *	Belger et al., 2011
25	Alaska Bethel, USA, NA	60°45'N, 161°45'W	Peatland <sup>a</sup>	<i>Empetrum nigrum, Carex aquatilis, Sphagnum spp.</i>	1988.7–1988.8	Bartlett et al., 1992;
26	Alaska Bethel, USA, NA	61°5'N, 162°1'W	Peatland <sup>a</sup>	<i>Empetrum nigrum, Carex aquatilis, Sphagnum spp.</i>	1988.7–1988.8	Fan et al., 1992
27	Alaska Prudhoe Bay, USA, NA	70°30'N, 149°00'W	Peatland <sup>a</sup>	<i>Sphagnum spp.</i>	1984 *	Sebacher et al., 1986
28	Alaska arboretum, USA, NA	64°52'N, 147°51'W	Peatland <sup>a</sup>	<i>Eriophorum vaginatum, Carex spp., Sphagnum spp.</i>	1987.6–1987.10 1988.6–1988.10 1989.6–1989.10	Whalen and Reeburgh, 1992
29	Saskatchewan, Canada, NA &	53°57'N, 105°57'W	Peatland	<i>Buckbean-Carex spp.</i>	1994.5–1994.9 1995.5–1995.10	Suyker et al., 1996; Sellers et al., 1997
30	Michigan, USA, NA	42°27'N, 84°01'W	Peatland	<i>Scheuchzeria palustris, Carex oligosperma</i>	1991.1–1993.12	Shannon et al., 1996
31	Toolik Lake, USA, NA &	68°38'N, 149°38'W	Peatland <sup>a</sup>	<i>Eriophorum, Carex spp.</i>	1990.6–1990.8 1992.6–1992.8 1993.5–1993.9	Christensen, 1993 Schimel et al., 1994, 1995
32	Hudson Bay, Canada, NA	51°18'–51°31'N, 80°28'–80°38'W	Peatland	<i>Larch, Black spruce, Sphagnum spp.</i>	1990.6–1990.10	Moore et al., 1994
33	Quebec, Canada, NA	54°48'N, 66°49'W	Peatland	<i>Carex spp.</i>	1989.6–1989.9	Moore et al., 1990
34	Mississippi, USA, NA	34°24'N, 89°50'W	Marsh	<i>Carex hyalinolepis, Hydrocotyle umbellata, Festuca obtusa</i>	2005.5–2006.7	Koh et al., 2009
35	Sherman Island, USA, NA	38°2'N, 121°45'W	Peatland	<i>Hordeum murinum L., Lepidium latifolium L.</i>	2009.4–2011.4 *	Hatala et al., 2012
36	Marcell forest, USA, NA	47°30'N, 93°29'W	Peatland	<i>Carex spp., sphagnum moss, Eriophorum chamissonis, etc.</i>	2009–2010 *	Olson et al., 2013
37	Mer Bleue peatland, Canada, NA	45°41'N, 75°48'W	Peatland	<i>Chamaedaphne calyculata, Ledum groenlandicum, etc.</i>	1999–2010 *	Moore et al., 2011
38	Sag river side, Alaska, NA	69°30'N, 148°13' W	Peatland <sup>a</sup>	<i>Vascular plant, moss and a few short shrubs</i>	1996.6–1996.9 *	Harazono et al., 2006
39	Happy valley, Alaska, NA	69°10'N, 148°51'W	Peatland <sup>a</sup>	<i>Sphagnum moss, sedge</i>	1995.6–1995.9 *	Harazono et al., 2006
40	Churchill Manitoba, Canada, NA	58°40'N, 93°50'W	Peatland	<i>Carex aquatilis, Eriophorum spp., etc.</i>	2008–2010 *	Hanis et al., 2013
41	Northern Alaska, USA, NA	71°17'N, 156°36'W	Peatland	<i>Moss, Carex aquatilis, Eriophorum vaginatum, etc.</i>	2007.6–2007.8 *	Zona et al., 2009
42	Alberta, Canada, NA	54°57'N, 112°28'W	Peatland	<i>Picea mariana, Larix laricina, shrub etc.</i>	2007.5–2007.9 *	Long et al., 2010
43	Great Dismal Swamp, USA, NA	35°54'N, 76°09'E	Swamp	<i>Taxodium distichum, Nyssa sylvatica, etc.</i>	2007.7–2009.6 *	Morse et al., 2012

\* We used the reported average yearly CH<sub>4</sub> flux of the experimental year/period from the literature.

& Wetland sites used for calibration.

<sup>§</sup> These swamp sites do not have plant species information in the literature.

<sup>a</sup> Tundra.

<sup>b</sup> Tidal marsh.

<sup>c</sup> Mangrove.

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**Table 2. Model performance for CH4MOD<sub>wetland</sub> and the TEM for different continents and wetland types**

Wetland type/Continent	CH4MOD <sub>wetland</sub>							TEM							n		
	R <sup>2</sup>	RMSE	RMD	EF	CD	U <sub>M</sub>	U <sub>R</sub>	U <sub>E</sub>	R <sup>2</sup>	RMSE	RMD	EF	CD	U <sub>M</sub>		U <sub>R</sub>	U <sub>E</sub>
North America	0.82	75.37	-1.96	0.57	0.49	0.04	0.61	0.39	0.80	56.22	-2.86	0.76	1.59	0.00	0.03	0.97	28
Asia	0.94	55.79	-12.64	0.93	0.96	0.28	0.02	0.70	0.26	72.56	1.71	0.32	1.93	0.00	0.03	0.97	11
Europe	0.35	62.69	-32.60	0.15	1.13	0.27	0.03	0.69	NS	161.33	29.39	-4.65	0.34	0.03	0.84	0.13	13
South America/Africa	NS	57.32	39.52	-0.80	0.67	0.48	0.07	0.45	0.59	29.33	13.13	0.53	2.22	0.13	0.04	0.83	6
Marsh	0.75	29.44	0.52	0.22	0.37	0.00	0.73	0.27	NS	39.76	-18.77	-0.42	0.95	0.22	0.17	0.61	8
Peatland	0.83	82.26	-10.4	0.57	0.49	0.02	0.61	0.38	0.70	69.45	7.96	0.69	1.14	0.01	0.03	0.96	39
Swamp	0.50	74.28	43.07	0.13	0.54	0.34	0.19	0.47	0.76	40.76	19.02	0.74	1.27	0.22	0.03	0.75	7
Coastal wetland	0.80	55.46	-26.97	0.72	2.09	0.24	0.30	0.47	NS	188.26	101.00	-2.26	0.35	0.29	0.42	0.29	4
Global	0.81	67.00	4.28	0.65	0.59	0.00	0.45	0.54	0.68	63.58	4.63	0.68	1.46	0.01	0.01	0.98	58

NS represents no significant correlation

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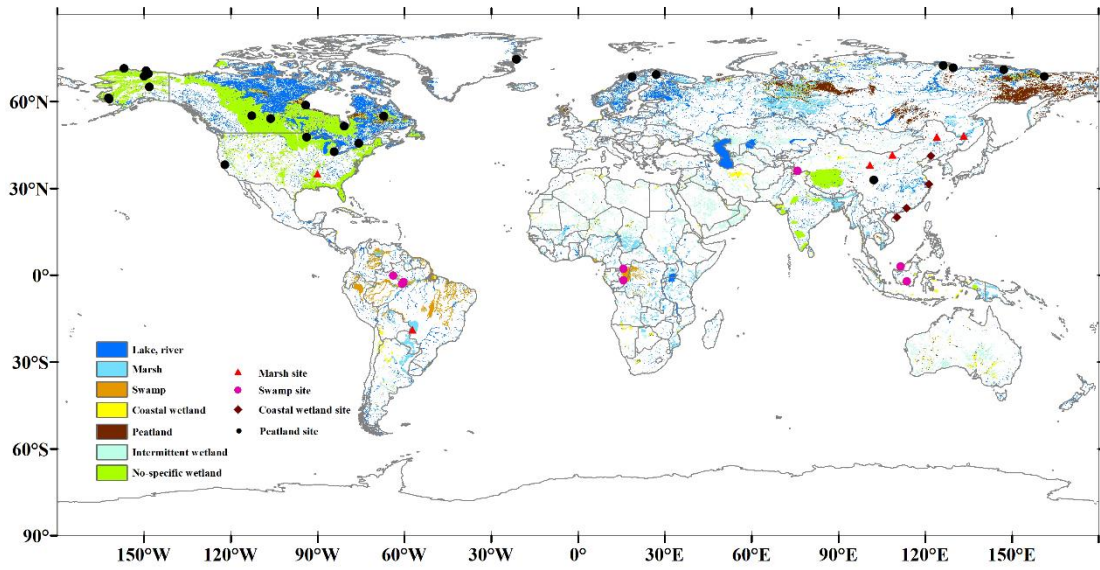
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903 **Table 3. CH<sub>4</sub> simulations by CH4MOD<sub>wetland</sub> and the TEM for different continents and wetland types. All units are Tg CH<sub>4</sub> yr<sup>-1</sup> ± 1σ, where the standard deviation represents the**  
 904 **interannual variation in the model estimates.**  
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Continent/Wetland type	CH <sub>4</sub> Flux (g m <sup>-2</sup> yr <sup>-1</sup> )		CH <sub>4</sub> Emissions (Tg)		Area (10 <sup>4</sup> × km <sup>2</sup> )
	CH4MOD <sub>wetland</sub>	TEM	CH4MOD <sub>wetland</sub>	TEM	
Asia	23.27 ± 0.67	25.78 ± 0.14	25.37 ± 0.73	28.81 ± 0.15	109.04
Africa	27.64 ± 1.55	33.92 ± 0.27	21.12 ± 1.18	25.91 ± 0.20	76.39
N. America	11.48 ± 0.47	14.10 ± 0.18	24.38 ± 1.00	29.95 ± 0.38	212.35
S. America	29.61 ± 0.52	39.91 ± 0.54	26.24 ± 0.46	35.36 ± 0.48	88.60
Europe	7.77 ± 0.11	16.04 ± 0.28	6.48 ± 0.09	13.38 ± 0.23	83.40
Oceania	12.08 ± 2.52	11.31 ± 0.57	1.72 ± 0.27	1.61 ± 0.06	14.23
Lake, river *	15.57	15.57	27.32	27.32	175.46
Marsh	28.64 ± 1.06	39.60 ± 0.28	37.47 ± 1.39	51.80 ± 0.37	131.61
Peatland	1.99 ± 0.09	9.87 ± 0.27	0.42 ± 0.02	2.06 ± 0.06	21.00
Swamp	31.58 ± 0.57	45.59 ± 0.91	17.37 ± 0.32	25.08 ± 0.50	55.34
Coastal wetland	9.44 ± 0.25	11.63 ± 0.14	2.85 ± 0.08	3.51 ± 0.04	30.32
Intermittent wetland	18.49 ± 2.08	18.29 ± 0.19	5.81 ± 0.65	5.75 ± 0.06	31.60
No-specific wetland	10.21 ± 0.68	13.63 ± 0.25	14.07 ± 0.93	18.79 ± 0.34	138.67
Global	18.03 ± 0.49	23.00 ± 0.15	105.31 ± 2.72	134.31 ± 0.84	584.00

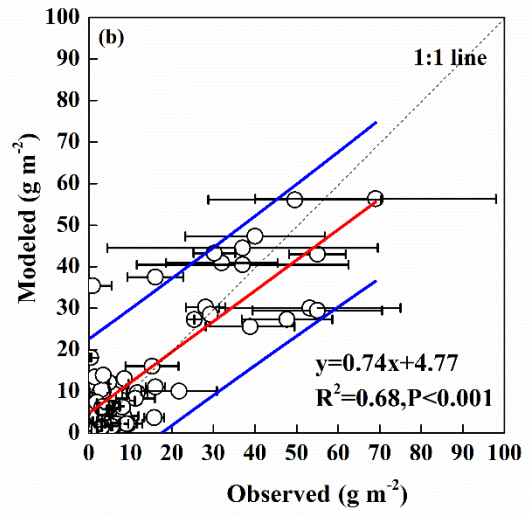
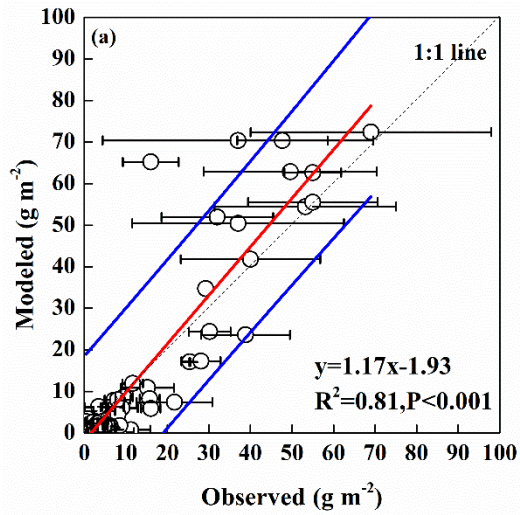
920 \* IPCC Tier 1 method was used to estimate the CH<sub>4</sub> emissions from lakes and rivers. The CH<sub>4</sub> emission factor was from the IPCC (1996).  
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 936 **Figure 1: Wetland site distribution (Table 1) and global wetland maps of GLWD-3 (Lehner and Döll, 2004).**

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Figure 2: Regression of simulated against observed total amount of seasonal/annual  $\text{CH}_4$  emissions from global wetland sites by  $\text{CH}_4\text{MOD}_{\text{wetland}}$  (a) and the TEM (b). The horizontal bars are the standard errors from the sampling replicates at the wetland site. The red line is the regression line of simulated vs. observed between modeled and observed values. The blue line is the prediction correspondence. The dashed line is the 1:1 line.

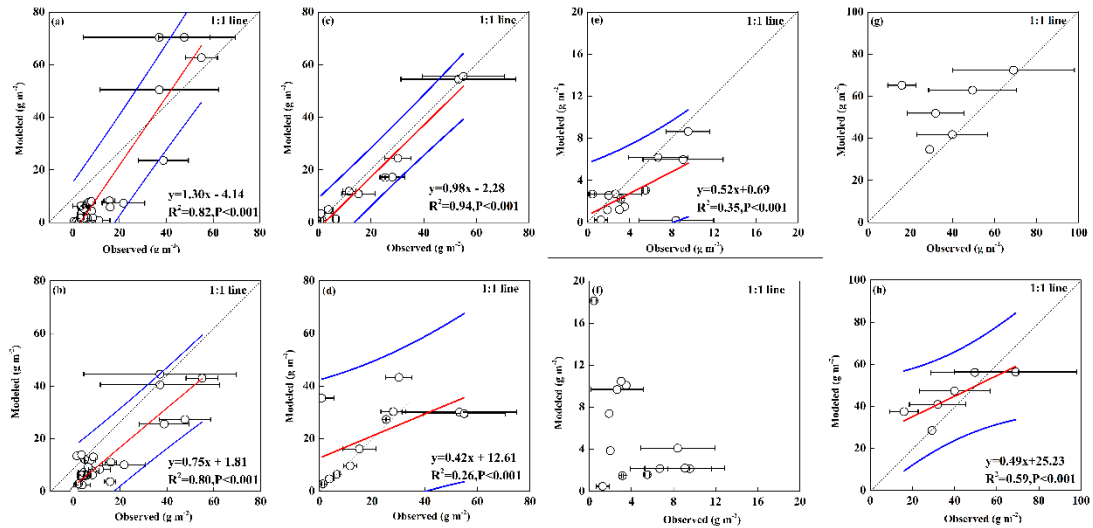
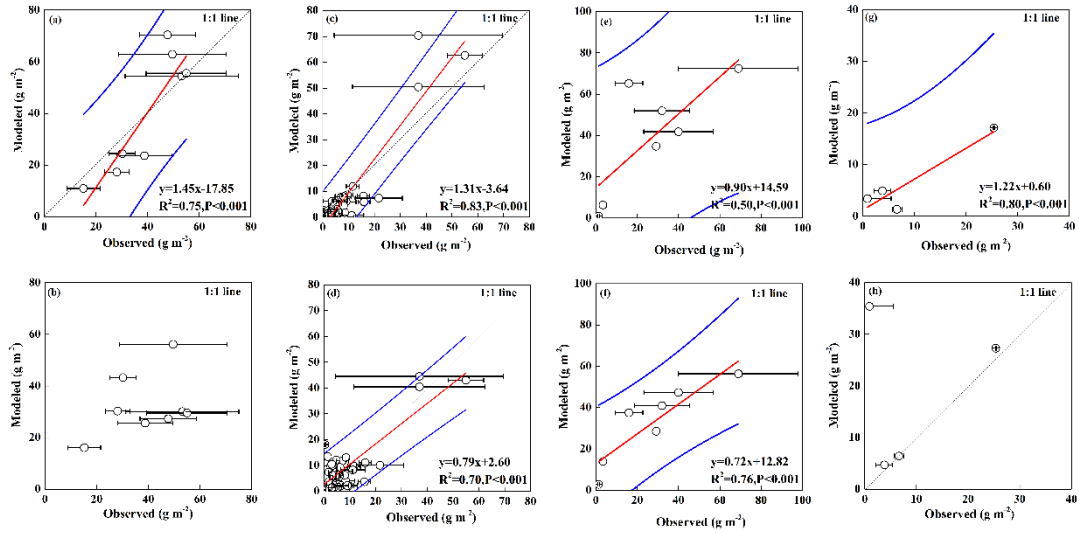


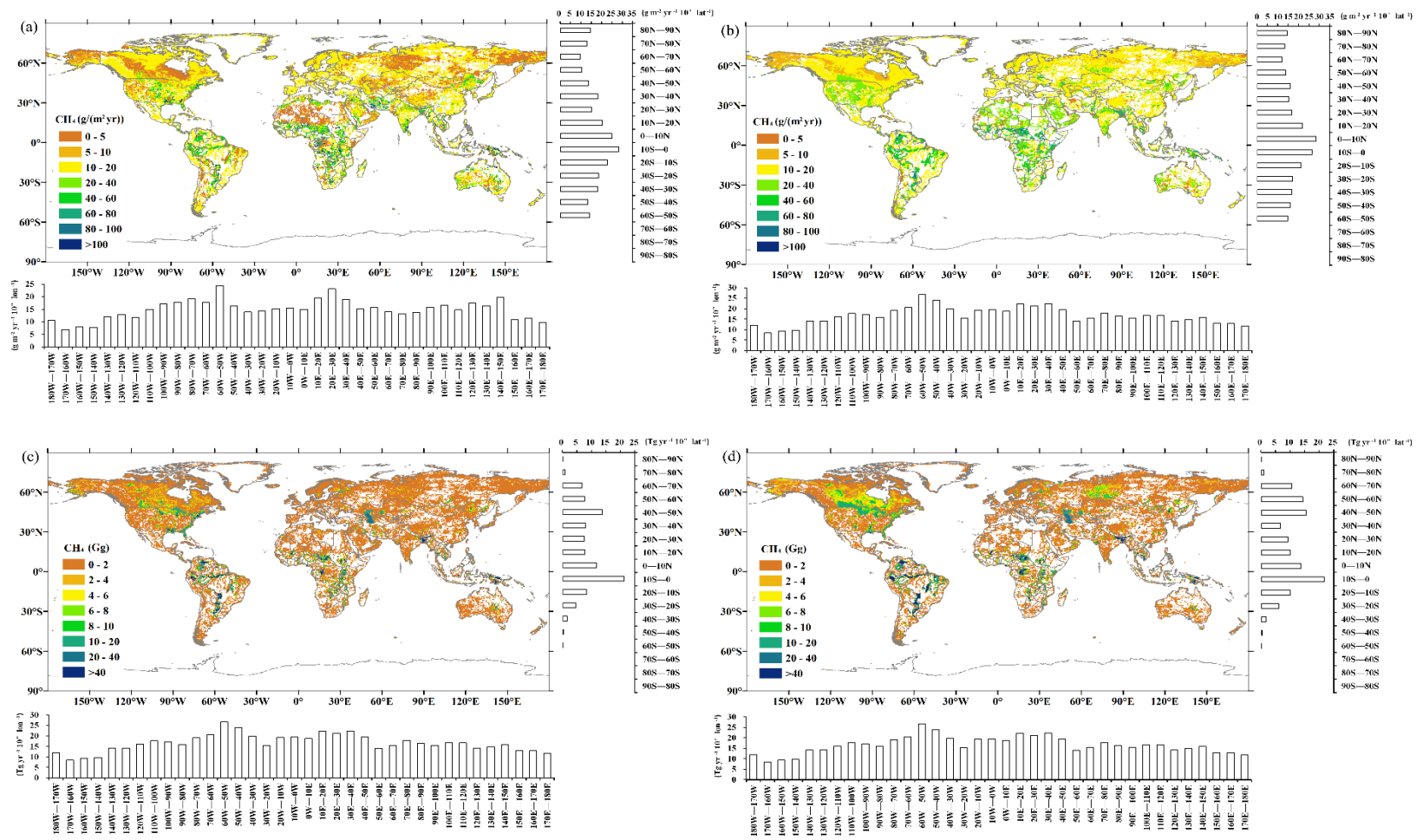
Figure 3: Regression of simulated against observed total amount of seasonal/annual CH<sub>4</sub> emissions from North American wetland sites by CH<sub>4</sub>MOD<sub>wetland</sub> (a) and the TEM (b), from Asian wetland sites by CH<sub>4</sub>MOD<sub>wetland</sub> (c) and the TEM (d), from European wetland sites by CH<sub>4</sub>MOD<sub>wetland</sub> (e) and the TEM (f), and from South American and African wetland sites by CH<sub>4</sub>MOD<sub>wetland</sub> (g) and the TEM (h). The horizontal bars are the standard errors from the sampling replicates at the wetland site. The blue line is the prediction correspondence. The dashed line is the 1:1 line.

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1001 **Figure 4: Regressions of simulated against observed total amount of seasonal/annual CH<sub>4</sub> emissions from**  
 1002 **marsh sites by CH<sub>4</sub>MOD<sub>wetland</sub> (a) and the TEM (b), from peatland sites by CH<sub>4</sub>MOD<sub>wetland</sub> (c) and the TEM**  
 1003 **(d), from swamp sites by CH<sub>4</sub>MOD<sub>wetland</sub> (e) and the TEM (f), and from coastal wetland sites by**  
 1004 **CH<sub>4</sub>MOD<sub>wetland</sub> (g) and the TEM (h). The horizontal bars are the standard errors from the sampling replicates**  
 1005 **at the wetland site. The blue line is the prediction correspondence. The dashed line is the 1:1 line.**

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Figure 5: Spatial pattern of annual mean CH<sub>4</sub> fluxes for 2000–2010, with latitudinal and longitudinal distributions of annual mean CH<sub>4</sub> fluxes by CH<sub>4</sub>MOD<sub>wetland</sub> (a) and the TEM (b). Spatial pattern of annual mean CH<sub>4</sub> emissions for 2000–2010, with latitudinal and longitudinal distributions of annual mean CH<sub>4</sub> emissions by CH<sub>4</sub>MOD<sub>wetland</sub> (c) and the TEM (d). The CH<sub>4</sub> fluxes and emissions are aggregated in steps of 10°.