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Simulating human impacts on global water resources using VIC-5

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11 Abstract. Questions related to historical and future water resources and water scarcity have been 12 addressed by several macro-scale hydrological models over the last few decades. However, further 13 advancements are needed to improve the integration of anthropogenic impacts and environmental flow 14 requirements into hydrological models. The newly developed VIC-WUR model aims to increase the 15 applicability of the VIC-5 model for water resource assessments, specifically by including human 16 impacts and environmental flow requirements. To this end, VIC-WUR extends VIC-5 with modules for 17 irrigation, domestic, industrial, energy and livestock water-use, environmental flow requirements for 18 surface and groundwater systems, and dam operation. Model inputs of sectoral water demand were 19 estimated independently and correlated well to reported national water withdrawals.

20 VIC-WUR results, based on the newly developed modules, corresponded with results from reported 21 global water withdrawals and other hydrological models, although differences exist. The VICWUR 22 irrigation withdrawals were high compared to the other models but closer to the reported values, 23 decreasing the gap between simulated and reported withdrawals. Irrigation withdrawals were probably 24 high due to the inclusion of groundwater withdrawals and paddy irrigation in the model. Domestic and 25 industrial water withdrawals were slightly lower than the reported values. Domestic and industrial 26 withdrawals were probably insufficient due to low water availability, as the potential water withdrawals 27 are more in line with reported values. Livestock water withdrawals were within the range of reported 28 values and other models.

The model additions comprehensively incorporate anthropogenic and environmental water use, which provides new opportunities for global water resource assessments. A preliminary assessment of environmental flow requirements shows competition between water resources allocated for human consumption and the environment, from ground and surface water sources. The improvements made here are a first step towards integrated water-food-energy nexus modelling.

34 1 Introduction

Questions related to historical and future water resources and scarcity have been addressed by several
 macro-scale hydrological models over the last few decades (Liang et al., 1994; Alcamo et al., 1997;





Hagemann and Gates, 2001; Takata et al., 2003; Krinner et al., 2005; Bondeau et al., 2007; Hanasaki et
al., 2008; Van Beek and Bierkens, 2008; Best et al., 2011). Early efforts focussed on the simulation of
natural water resources and the impacts of land cover and climate change on water availability (Oki et
al., 1995; Nijssen et al., 2001a; Nijssen et al., 2001b). Recently, a larger focus has been on incorporating
anthropogenic impacts, such as water withdrawals and dam operations, into water resource assessments
(Alcamo et al., 2003; Haddeland et al., 2006b; Biemans et al., 2011; Wada et al., 2011b; Hanasaki et al.,
2018).

44 Global water withdrawals increased eight-fold over the last century and are projected to increase further 45 (Shiklomanov, 2000; Wada et al., 2011a). Although water withdrawals are only a small fraction of the total global runoff (Oki and Kanae, 2006), water scarcity can be severe due to the variability of water in 46 47 both time and space (Postel et al., 1996). Already severe water scarcity is experienced by two-thirds of the global population for at least part of the year (Mekonnen and Hoekstra, 2016). To stabilize water 48 49 availability for different sectors (e.g. irrigation, hydropower, and domestic uses) dams and reservoirs 50 were built, which are able to strongly affect global river discharge (Nilsson et al., 2005; Grill et al., 51 2019). In addition, groundwater resources are being extensively exploited to meet increasing water 52 demands (Rodell et al., 2009; Famiglietti, 2014).

53 However, further advancements are needed to improve the integration of anthropogenic impacts into 54 hydrological models (Döll et al., 2016). Several models do not yet incorporate all aspects of 55 anthropogenic water withdrawals such as domestic, manufacturing and energy (thermoelectric) water 56 withdrawals from both ground and surface water. Although these sectors use less water than irrigation 57 (Shiklomanov, 2000; Grobicki et al., 2005; Hejazi et al., 2014) they are locally important actors (Gleick 58 et al., 2013), especially for the water-food-energy nexus (Bazilian et al., 2011). Sufficient water supply 59 and availability are essential for meeting a range of local and global sustainable development goals 60 related to water, food, energy and ecosystems (Bijl et al., 2018).

Environmental flow requirements (EFRs) are also often neglected in global water resource assessments
(Pastor et al., 2014), even though they are "(...) necessary to sustain aquatic ecosystems which, in turn,
support human cultures, economies, sustainable livelihoods, and well-being" (Brisbane Declaration,





- 64 2017). Various EFR methods are available for streamflow (Smakhtin et al., 2004; Richter et al., 2012; 65 Pastor et al., 2014) and groundwater (Gleeson and Richter, 2018), although environmental limits for 66 groundwater withdrawal have only recently been considered explicitly. Anthropogenic alterations 67 already strongly affect freshwater ecosystems (Carpenter et al., 2011), with more than a quarter of all 68 global rivers experiencing very high biodiversity threats (Vorosmarty et al., 2010). By neglecting EFRs, 69 water availability for anthropogenic uses is likely over-estimated (Gerten et al., 2013).
- 70 One of widely-used macro-scale hydrological models is the Variable Infiltration Capacity (VIC) model. 71 The model was originally developed as a land-surface model (Liang et al., 1994), but has been mostly 72 used as a stand-alone hydrological model (Abdulla et al., 1996; Nijssen et al., 1997) using an offline 73 routing module (Lohmann et al., 1996; Lohmann et al., 1998b, a). Where land-surface models focus on 74 the vertical exchange of water and energy between the land surface and the atmosphere, hydrological 75 models focus on the lateral movement and availability of water. By combining these two approaches, 76 VIC simulations are strongly process-based and this, in turn, provides a good basis for climate-impact 77 modelling. Recently version 5 of the VIC model (VIC-5) was released (Hamman et al., 2018), which 78 focussed on improving the model infrastructure. These improvements are highly relevant when 79 simulating anthropogenic impacts on global water resources.

80 VIC has been used extensively in studies ranging from: coupled regional climate model simulations 81 (Zhu et al., 2009; Hamman et al., 2016), combined river discharge and water-temperature simulations 82 (van Vliet et al., 2016), hydrological sensitivity to climate change (Hamlet and Lettenmaier, 1999; 83 Nijssen et al., 2001a; Chegwidden et al., 2019), global streamflow simulations (Nijssen et al., 2001b), 84 and real-time drought forecasting (Wood and Lettenmaier, 2006; Mo, 2008). Several studies used VIC 85 to simulate the anthropogenic impacts of irrigation and dam operation on water resources (Haddeland et al., 2006a; Haddeland et al., 2006b; Zhou et al., 2015; Zhou et al., 2016) based on the model setup of 86 87 Haddeland et al. (2006b). However, water withdrawals for other sectors and flow requirements for 88 freshwater ecosystems were ignored in these studies.

Our study aims to increase the applicability of the VIC-5 model for water resource assessments,
 specifically by including human impacts and environmental flow requirements. Here the newly





- 91 developed VIC-WUR model is presented (named after the developing team at Wageningen University 92 and Research). The VIC-WUR model extends the existing VIC-5 model with several modules that 93 simulate the anthropogenic impacts on water resources. These modules include: integrated routing, 94 water use for various sectors (irrigation, domestic, industrial, energy and livestock), environmental flow 95 requirements for both surface and groundwater systems, and dam operation.
- 96 The next section first describes the original VIC-5 hydrological model (Section 2.1), which calculates 97 natural water resource availability. Subsequently the integration of the anthropogenic impact modules, 98 which modify the water resource availability, are described (Section 2.2). Global anthropogenic water 99 uses for each sector are also estimated (Section 2.3). To assess the performance of the newly developed modules, the VIC-WUR results were compared with reported global withdrawal data from Shiklomanov 100 101 (2000) and Steinfeld et al. (2006) as well as various other state-of-the-art global hydrological models 102 used in the Inter-Sectoral Impact Model Intercomparison Project (ISIMIP; Warszawski et al., 2014) and 103 Water and Global Change project (WATCH; Harding et al., 2011) (Section 3.1). The results also contain 104 a preliminary assessment of the water availability constrains imposed by EFRs.

105 2 Model development

106 2.1 VIC hydrological model

107 The basis of the VIC-WUR model is the Variable Infiltration Capacity model version 5 (VIC-5) (Liang 108 et al., 1994; Hamman et al., 2018). VIC-5 is an open source macro-scale hydrological model that 109 simulates the full water and energy balance on a (latitude - longitude) grid. Each grid cell accounts for 110 sub-grid variability in land cover and topography, and allows for variable saturation across the grid cell. 111 For each sub-grid the water and energy balance is computed individually (i.e. sub-grid do not exchange 112 water or energy between one another). The methods used to calculate the water and energy balance are 113 summarized in Appendix A, mainly based on the work of Liang et al. (1994). For the description of the 114 global calibration and validation of the water balance one is referred to Nijssen et al. (2001b).

VIC version 5 (Hamman et al., 2018) upgrades did not change the model representation of physical
processes, but improved the model infrastructure. Improvements include the use of NetCDF for





- 117 input/output and the implementation of parallelization through Message Passing Interface (MPI). These changes increase computational speed and make VIC-5 better suited for (computationally expensive) 118 119 global simulations. The most significant modification that enables new model applications is that VIC-120 5 also changed the processing order of the model. In previous versions all timesteps were processed for 121 a single grid cell before continuing to the next cell (time-before-space). In VIC-5 all grid cells are 122 processed before continuing to the next timestep (space-before-time). This development allows for 123 interaction between grid cells every timestep, which is important for full integration of the anthropogenic 124 impact modules, especially water withdrawals and dam operation.
- For example, surface and subsurface runoff routing to produce river streamflow was typically done as a post-process operation (Lohmann et al., 1996; Hamman et al., 2017), due to the time-before-space processing order of previous versions. Therefore, water withdrawals could not be taken into account directly and studies using the model setup of Haddeland et al. (2006b) required multiple successive model runs. Since VIC-5 uses the space-before-time processing order, runoff routing could be simulated each timestep. The routing post-process was replaced by our newly developed routing module which simulates routing sequentially (upstream-to-downstream) to facilitate water withdrawals between cells.
- 132 2.2 Anthropogenic-impact modules

VIC-WUR extends the existing VIC-5 though the addition of several newly developed anthropogenicimpact modules (Figure 1). These modules include sector-specific water withdrawal and consumptions,
environmental flow requirements for both surface and groundwater systems and dam operation for small
(within-grid) and large dams.







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Figure 1: Schematic overview of the VIC-WUR model that includes the VIC-5 model and several anthropogenic impact modules. Water from river streamflow, groundwater and small (within-grid) reservoirs are available for withdrawal. Surface and groundwater withdrawals are constrained by environmental flow requirements. Withdrawn water is available for irrigation, domestic, industrial, energy and livestock use. Unconsumed irrigation water is returned to the soil column of the hydrological model. Unconsumed water for the other sectors is returned to the river streamflow. Small reservoirs fill using surface runoff from the cell they are located, while large dam reservoirs operate solely on rivers streamflow.

145 2.2.1 Water withdrawal and consumption

In VIC-WUR, sectoral water demands need to be specified for each grid cell (Section 2.3). To meet 146 147 water demands, water can be withdrawn from river streamflow, small (within-grid) reservoirs and 148 groundwater resources. Streamflow withdrawals are abstracted from the grid cell discharge (as generated by the routing module), reservoir withdrawals are abstracted from small dam reservoirs 149 150 (located in the cell) and groundwater withdrawals are abstracted from the third layer soil moisture. The 151 partitioning of water withdrawals between surface and ground water resources was based on the study 152 of Döll et al. (2012), who estimated the groundwater withdrawal fraction for each sector in around 153 15.000 national and sub-national administrative units. Groundwater fractions were based mainly on 154 information from the International Groundwater Resources Assessment Centre (IGRAC; un-igrac.org) 155 database. Surface water withdrawals are partitioned between river streamflow and small reservoirs 156 relative to water availability.

Water could also be withdrawn from the river streamflow of other 'remote' cells in delta areas. Since rivers cannot split in our routing module, the model is unable to simulate the redistribution of water





- resources in dendritic deltas. Therefore, streamflow at the river mouth is available for use in delta areas
 to simulate the actual water availability. Delta areas were delineated by the global delta map of Tessler
 et al. (2015).
- 162 When water demands cannot be met, water withdrawals are allocated to the domestic, energy, 163 manufacturing, livestock and irrigation sector in that order. Withdrawn water is partly consumed, 164 meaning the water evaporates and does not return to the hydrological model. Consumption rates were 165 set at 0.15 and 0.10 for the domestic and industrial sectors respectively, based on the data of 166 Shiklomanov (2000). The water consumption in the energy sector was based on Goldstein and Smith 167 (2002) and varies per thermoelectric plant based on the fuel type and cooling system. For the livestock 168 sector the assumption was made that all withdrawn water is consumed. Unconsumed water withdrawals 169 for these sectors are returned as river streamflow. For the irrigation sector, consumption was determined 170 by the calculated evapotranspiration. Unconsumed irrigation water remains in the soil column and 171 eventually returns as subsurface runoff.

172 2.2.2 Environmental flow requirements

Water withdrawals can be constrained by environmental flow requirements (EFRs). These EFRs specify the timing and quantity of water needed to support terrestrial river ecosystems (Smakhtin et al., 2004; Pastor et al., 2019). Surface and groundwater withdrawals are constrained separately in VIC-WUR, based on the EFRs for streamflow and baseflow respectively. EFRs for streamflow specify the minimum river streamflow requirements while EFRs for baseflow specify the minimum subsurface runoff requirements (from groundwater to surface water). Since baseflow is a function groundwater availability in the hydrological model, baseflow requirements are used to constrain groundwater withdrawals.

Our study used the Variable Monthly Flow (VMF) method (Pastor et al., 2014) to calculate the EFRs for streamflows. VMF calculates the required streamflow as a fraction of the natural flow during high (30%), intermediate (45%) and low (60%) flow periods, as described in Appendix B. The VMF method performed favourably compared to other hydrological methods, such as the method proposed by Smakhtin et al. (2006) or the Q90-Q50 method, in 11 case studies where EFRs were calculated locally





- 185 (Pastor et al., 2014). The advantage of the VMF method is that the method accounts for the natural flow
- 186 variability, which is essential to support freshwater ecosystems (Poff et al., 2010).
- 187 EFR methods for baseflow have been rather underdeveloped compared to EFR methods for streamflow. 188 However, a presumptive standard of 90 % of the natural subsurface runoff through time was proposed 189 by Gleeson and Richter (2018), as described in Appendix B. This standard should provide high levels 190 of ecological protection, especially for groundwater dependent ecosystems. Note that part of the EFRs 191 for baseflow are already captured in the EFRs for streamflow, especially during low-flow periods that 192 are usually dominated by baseflows. However, the EFRs for baseflow specifically limit local 193 groundwater withdrawals while EFRs for streamflow include the accumulated runoff from upstream 194 areas. Also, the chemical composition of groundwater derived flows is inherently different, making them 195 a non-substitutable water flow for environmental purposes (Gleeson and Richter, 2018).

196 2.2.3 Dam operation

197 Due to the lack of globally available information on local dam operations, several generic dam operation 198 schemes were developed for macro-scale hydrological models to reproduce the effect of dams on natural 199 streamflow (Haddeland et al., 2006a; Hanasaki et al., 2006; Zhao et al., 2016). In VIC-WUR a 200 distinction is made between 'small' dam reservoirs (with an upstream area smaller than the cell area) 201 and 'large' dam reservoirs, similar to Hanasaki et al. (2018), Wisser et al. (2010) and Döll et al. (2009). 202 Small dam reservoirs act as buckets that fill using surface runoff of the grid-cell they are located in and 203 reservoirs storage can be used for water withdrawals in the same cell. Large dam reservoirs are located 204 in the main river and used the operation scheme of Hanasaki et al. (2006).

The scheme distinguishes between two dam types: (1) dams that do not account for water demands downstream (e.g. hydropower dams or flood protection dams) and (2) dams that do account for water demand downstream (e.g. irrigation dams). For dams that do not account for demands, dam release is aimed at reducing annual fluctuations in discharge. For dams that do account for demands, dam release is additionally adjusted to provide more water during periods of high demand. The operation scheme was validated by Hanasaki et al. (2006) for 28 reservoirs and was used in various studies (Hanasaki et





- 211 al., 2008; Döll et al., 2009; Pokhrel et al., 2012; Voisin et al., 2013; Hanasaki et al., 2018). Here, the
- 212 scheme was adjusted slightly to account for monthly varying EFRs and to reduce overflow releases. The
- 213 full operation scheme is described in Appendix C.
- 214 The Global Reservoir and Dam (GRanD) database (Lehner et al., 2011) was used to specify location,
- 215 capacity, function (purpose), and construction year of each dam. The capacity of multiple (small- and
- 216 large) dams located in the same cell were combined.

217 2.3 Sectoral water demands

218 Water withdrawals are based on the irrigation, domestic, industry, energy and livestock water demand 219 in each grid-cell. Water demands represent the potential water withdrawal, which is reduced when 220 insufficient water is available. Irrigation demands were estimated based on the hydrological model while 221 water demands for other sectors are provided to the model as an input. Domestic and industrial were 222 estimated based on several socioeconomic predictors, while energy and livestock water demands were 223 mostly data-driven (i.e. derived from power plant and livestock distribution data). Due to data limitations 224 the energy sector was incomplete, and energy water demands were partly included in the industrial water 225 demands (which combined the remaining energy and manufacturing water demands). For more details 226 concerning sectoral water demand calculations the reader is referred to Appendix D.

227 2.3.1 Irrigation demands

228 Irrigation demands were set to increase soil moisture in the root zone so that water availability is not 229 limiting crop evapotranspiration and growth. Preferably, irrigation was supplied to fill the soil to field 230 capacity (Allen et al., 1998), which is the moisture content where water leaching is minimized. The 231 exception is paddy rice irrigation (Brouwer et al., 1989), where irrigation was also supplied to keep the 232 upper soil layer saturated. Water demands for paddy irrigation practices are relatively high compared to 233 conventional irrigation practices due to increased evaporation and percolation. Therefore, the crop 234 irrigation demands for these two irrigation practices were calculated and applied separately (i.e. in 235 different sub-grids).





Total irrigation demands also included transportation and application losses. Note that transportation and application losses are not 'lost' but rather returned to the soil column without being used by the crop. The water loss fraction was based on Frenken and Gillet (2012), who estimated the irrigation efficiency for 22 United Nations sub-regions based on differences between calculated irrigation requirements and reported irrigation withdrawals. Potential total irrigation demands were validated independently and correlated well with reported withdrawals (adjusted $R^2 > 0.8$; Figure 2a).

242 2.3.2 Domestic and industrial demands

243 Domestic and industrial water withdrawals were estimated based on Gross Domestic Product (GDP) per 244 capita and Gross Value Added (GVA) by industries respectively. These drivers do not fully capture the 245 multitude of socioeconomic factors that influence water demands (Babel et al., 2007). However, the 246 wide availability of data allows for extrapolation of water demands to data-scarce regions and future 247 scenarios (using studies such as Chateau et al. (2014)).

248 Domestic water demands per capita (used for drinking, sanitation, hygiene and amenity uses) were 249 estimated similar to Alcamo et al. (2003). Demands increased non-linearly with GDP per capita due to 250 the acquisition of water using appliances as household become richer. A minimum water supply is 251 needed for survival, and the saturation of water using appliances sets a maximum on domestic water 252 demands. Industrial water demands (used for cooling, transportation and manufacturing) were estimated 253 similar to Flörke et al. (2013) and Voß and Flörke (2010). Industrial demands increased linearly with 254 GVA (as an indicator of industrial production). Since industrial water intensities (i.e. the water use per 255 production unit) vary widely between different industries (Flörke and Alcamo, 2004; Vassolo and Döll, 256 2005; Voß and Flörke, 2010), the average water intensity was estimated for each country. Both domestic 257 and industrial water demands were also influenced by technological developments that increase water-258 use efficiency over time, as in Flörke et al. (2013). Estimated domestic and industrial water demands 259 were validated independently and correlated well to reported withdrawals (adjusted R² > 0.8; Figure 2b 260 and Figure 2c).





261 Domestic water demands varied monthly based on air temperature variability as in Wada et al. (2011b). 262 Using this approach, water demands were higher in summer than in winter, especially for counties with 263 strong seasonal temperature differences. Domestic water demand per capita were downscaled using the 264 HYDE3.2 gridded population maps (Goldewijk et al., 2017). Industrial water demands were kept 265 constant throughout the year. Industrial demands were downscaled from national to grid cell values 266 using the NASA Back Marble night-time light intensity map (Roman et al., 2018). National industrial 267 water demands were allocated based on the relative light intensity per grid cell for each country.

268 2.3.3 Energy and livestock demands

269 Energy water demands (used for cooling of thermoelectric plants) were estimated using data from van 270 Vliet et al. (2016). Water use intensity for generation (i.e. the water use per generation unit) was 271 estimated based on the fuel and cooling system type (Goldstein and Smith, 2002), which was combined 272 with the installed generation capacity. Note that the data only covered a selection of the total number of 273 thermoelectric power plants worldwide. Around 27% of the total (non-renewable) global installed 274 capacity between 1980 and 2011 was included in this dataset due to lack of information on cooling 275 system types for the majority of thermoelectric plants. To avoid double counting, energy water demands 276 were subtracted from the industrial water demands.

Livestock water demands (used for drinking and animal servicing) were estimated by combining the Gridded Livestock of the World (GLW3) map (Gilbert et al., 2018) with the livestock water requirement reported by Steinfeld et al. (2006). Eight varieties of livestock were considered: cattle, buffaloes, horses, sheep, goats, pigs, chicken and ducks. Drinking water demands varied monthly based on temperature as described by Steinfeld et al. (2006), whereby drinking water requirements were higher during higher temperatures.

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Figure 2: Comparison between reported and estimated national water withdrawal per year for the irrigation (a), domestic (b) and industrial (c) sector. Reported values are from the validation dataset. Note the log-log axis which is used to display the wide range of water withdrawals. The adjusted R squared is also based on the log-log values.

- 288 **3** Model application
- 289 3.1 Setup

VIC-WUR results were generated between 1979 and 2016, excluding a spin-up period of one year (analysis period from 1980 to 2016). The model used a 6-hourly timestep and simulations were executed on a 0.5° by 0.5° grid (around 55 km at the equator) with three soil layers per grid cell. Soil and (natural) vegetation parameters were the same as in Nijssen et al. (2001c) (disaggregated to 0.5°), who used various sources to determine the soil (Cosby et al., 1984; Carter and Scholes, 1999) and vegetation parameters (Calder, 1993; Ducoudre et al., 1993; Sellers et al., 1994; Myneni et al., 1997).

296 Nijssen et al. (2001c) used the Advanced Very High Resolution Radiometer vegetation type database 297 (Hansen et al., 2000) to spatially distinguish 13 land cover types. The land cover type 'cropland' in the 298 original land-cover dataset was replaced by cropland extents from the MIRCA2000 cropland dataset 299 (Portmann et al., 2010). MIRCA2000 distinguishes the monthly growing area(s) and season(s) of 26 300 irrigated and rain-fed crop types around the year 2000. Crop types were aggregated into three land cover 301 types: rain-fed, irrigated and paddy rice cropland. The natural vegetation was proportionally rescaled to 302 make up discrepancies between the natural vegetation and cropland extents. 303 Cropland coverage (the cropland area actually growing crops) varied monthly based on the crop growing

304 areas of MIRCA2000. The remainder was treated as bare soil. Cropland vegetation parameters (e.g. Leaf





- Area Index (LAI), displacement, vegetation roughness and albedo) vary based on the monthly crop growing seasons and the development-stage crop coefficients of the Food and Agricultural Organisation
- 307 (Allen et al., 1998).
- The latest WATCH forcing data Era Interim (aggregated to 6 hourly), developed by the EU Water and Global Change (WATCH; Harding et al., 2011) project, was used as climate forcing (WFDEI; Weedon et al., 2014). The dataset provides gridded historical climatic variables of minimum and maximum air temperature, precipitation (as the sum of snowfall and rainfall, GPCC bias-corrected), relative humidity, pressure and incoming shortwave and longwave radiation.

313 3.2 Results

314 The VIC-WUR model results were compared to several of the Inter-Sectoral Impact Model 315 Intercomparison Project (Warszawski et al., 2014) simulation round 2a global hydrological impact 316 models: H08 (Hanasaki et al., 2008), LPJmL (Sitch et al., 2003), VIC (Liang et al., 1994), PCR-317 GLOBWB (Wada et al., 2014) and WaterGAP (Muller Schmied et al., 2016). The ISIMIP2a outputs are 318 comparable to the results of our study since the same meteorological and land cover inputs were used. 319 The VIC and LPJmL models only provided data on the actual and potential irrigation withdrawal and 320 consumption. H08 additionally provided data for the domestic sector, and PCR-GLOBWB additionally 321 provided data for the domestic and livestock sector. To increase the number of models to compare to, 322 the WaterGAP (Alcamo et al., 2003) output for the domestic and industrial (manufacturing plus energy) 323 sector from the Water and Global Change project (Harding et al., 2011) was included as well. Note that 324 the WaterGAP simulations were based on a different WATCH forcing dataset (WFD) (Weedon et al., 325 2011). Since our study used a present-day land-cover map, the actual and potential irrigation 326 withdrawals were compared based on the so-called 'pressoc' (present-day land-cover) simulations. 327 Domestic, industrial and livestock sectors were compared based on the 'varsoc' (variable human 328 influences) simulations, since our study used varying socioeconomic predictors to estimate the water 329 demand in these sectors. Results were compared between the years 1980 to 2005. The reported global 330 water withdrawal of Shiklomanov (2000) and Steinfeld et al. (2006) are included as a reference.





331 3.2.1 Irrigation sector

332 Compared to other models the VIC-WUR potential and actual water withdrawals (without EFRs) were 333 at the high end (Figure 3). Annual potential and actual irrigation withdrawals for VIC-WUR were around 334 3060 km³ and 1870 km³ respectively, while the ensemble mean potential and actual withdrawals were 335 only 2200 km³ and 1400 km³ respectively. Especially in the African and Asian regions the irrigation 336 withdrawals were high compared to the model ensemble. Irrigation withdrawals were probably high due 337 to the inclusion of groundwater withdrawals and paddy irrigation in the model. All models (VIC-WUR 338 included) indicated a lower actual irrigation withdrawal than reported. Due to the increased irrigation 339 withdrawal, the deficit for VIC-WUR (around 710 km³) was lower than the ensemble mean deficit 340 (around 1170 km³). This difference is often assumed to be met by non-renewable and/or unspecified 341 withdrawals (Wada et al., 2010; Hanasaki et al., 2018).



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Figure 3: Irrigation withdrawals for the world (a) and each region (b-f) for five hydrological models between 1980 and 2005. Colours differentiate between the potential (green), actual excluding EFRs (orange) and actual including EFRs (purple) withdrawals. The spread indicates the inter-annual variation of the simulated irrigation withdrawals. Data was obtained from the ISIMIP project, except for the VIC-WUR model. The black line indicates the reported total irrigation withdrawal as estimated by Shiklomanov (2000). Note that the y-axis varies for each graph, making Asia is by far the largest contributor to irrigation withdrawals.

349 Limitations imposed by the environmental flow requirements reduced the actual (irrigation) water

350 withdrawals by about 43% (Figure 4a). In total, 71% of the reduction could be attributed to limitations





- 351 imposed on groundwater withdrawals. Therefore, the impact of the environmental flow requirements
- 352 was largest in groundwater dependent regions (Figure 4b). However, surface water withdrawals
- increased by 11 % when limiting groundwater withdrawals on top of limiting surface water withdrawals,
- due to subsurface runoff increases.



Figure 4: Average annual water withdrawals reductions when adhering to EFRs as (a) global gross total and (b) spatially distributed. Global gross totals (a) are separated into withdrawals without EFRs (red), withdrawals with EFRs for streamflow (purple) and withdrawals with EFRs for both streamflow and baseflow (blue). Note the log axis for the spatially distributed withdrawal reductions (b) to better display the spatial distribution of the reductions. Blue regions indicate areas where the withdrawal reduction is largely (> 75 %) caused by the EFRs for baseflow.

361 3.2.2 Domestic, industrial and livestock sector

In contrast to the irrigation withdrawals, the annual domestic (ranging from 195 km³ to 275 km³) and 362 industrial (ranging from 461 km³ to 637 km³) withdrawals of VIC-WUR were slightly lower than that 363 of reported values and other models (Figure 5a; Figure 5b). Domestic and industrial withdrawals were 364 probably low due to insufficient water availability, as the potential water withdrawals are more in line 365 366 with reported values. Note that the rising trend of the VIC-WUR domestic water withdrawals (on average 2.5 km³ year⁻¹), WaterGAP (on average 2.3 km³ year⁻¹) and H08 (on average 2.4 km³ year⁻¹) 367 was more gradual than that of PCR-GLOBWB (on average 8.3 km³ year⁻¹) and Shiklomanov (2000) (on 368 average 8.1 km³ year⁻¹) between 1980 and 2000. This slope difference resulted from the different 369





- 370 methods used to calculate domestic water demand. The annual livestock withdrawals of VIC-WUR
- 371 (ranging from 25 km³ to 27 km³) were within range of reported values and other models (Figure 5c).



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Figure 5: Domestic (a), industrial (b) and livestock (c) water withdrawals for the world and each region for five hydrological models between 1980 and 2005. Colours differentiate between models. Data was obtained from the ISIMIP project (for H08 and PCR-GLBOWB) and WATCH project (for WaterGAP), except for the VIC-WUR model. The black points indicates the reported total water withdrawals for each sector as estimated by Shiklomanov (2000) (for domestic and industrial sectors) and (Steinfeld et al., 2006) (for livestock sector). Note that the y-axis varies for each graph, making Asia, Europe and Asia the largest contributors to the domestic, industrial and livestock water withdrawals respectively.

380 4 Discussion

- 381 Our paper presents the newly developed VIC-WUR model that aims to provide new opportunities for
- 382 global water resource assessments by integrating several anthropogenic-impact modules. The results of
- the VIC-WUR model are in line with reported water withdrawal values of Shiklomanov (2000) and





384 (Steinfeld et al., 2006), as well as the results of other hydrological models available via the ISIMIP and

385 WATCH projects. However, there are some important differences.

Potential irrigation withdrawal differences between models reflect the differences in the representation 386 387 of hydrological processes as well as the method used to calculate irrigation demands. Especially the 388 differences between VIC and VIC-WUR are interesting since they both employ the same hydrological 389 model. The increase in potential irrigation withdrawal between VIC and VIC-WUR can be attributed to 390 the inclusion of paddy irrigation by VIC-WUR. VIC irrigates crops only when they experience water 391 stress, while in VIC-WUR paddy irrigation is also used to saturate the top soil layer. The VIC-WUR potential withdrawals are 56 % higher than the VIC potential withdrawals for cells where rice is the 392 major crop (> 50 % of cropland). Potential irrigation withdrawals for convention irrigation is actually 393 394 higher for VIC than VIC-WUR, since the field capacity in VIC-WUR is tuned lower than that the field 395 capacity in VIC (see Appendix D). The lower field capacity results in reduced percolation for conventional irrigation. The VIC potential withdrawals are 33 % higher than the VIC-WUR potential 396 397 withdrawals for cells where no rice is present. This difference is reflected in the spatial distribution of 398 water demands. Potential irrigation withdrawals of VIC-WUR are higher than those of VIC, except for 399 the Americas and Europe where paddy irrigation is relatively limited.

400 Actual irrigation withdrawals of VIC-WUR are high compared to the other models. This difference can 401 be explained, in part, since some models (LPJmL, VIC and H08) did not (yet) include groundwater 402 withdrawals in their simulations. The high irrigation withdrawals of VIC-WUR decrease the gap 403 between the reported and simulated irrigation withdrawals often assumed to be met by (non-renewable) 404 groundwater withdrawals, in particular fossil groundwater stores (Wada et al., 2010). Often the regions 405 where the simulated withdrawals are lower than the actual withdrawals are regions where unsustainable 406 groundwater exploitation is reported by several studies (Gleeson et al., 2012; Rodell et al., 2018). To 407 our knowledge no previous study has estimated the amount of global non-renewable groundwater 408 withdrawals without using one of the models mentioned above. Therefore, the accuracy of actual 409 irrigation withdrawal results cannot be verified.





- 410 Differences in domestic, industrial and livestock actual water withdrawals between models are difficult 411 to explain since most studies use different methods to calculate and downscale sectoral water demands. 412 Therefore, water availability is not the only factor affecting the actual water withdrawals. Inputs of 413 potential domestic and industrial water withdrawals are close to the values reported by Shiklomanov 414 (2000). However, the actual water withdrawals are lower, indicating limited water availability. The lack 415 of water availability could be due to a number of factors: (1) The spatial distribution of water demands, 416 (2) the division between groundwater and surface water withdrawals, and/or (3) simulations of water 417 availability are insufficient in certain regions. Improvements would require more data to improve 418 groundwater and surface water demands and/or regional verification of water availability.
- 419 Environmental Flow Requirements (EFRs) for both baseflow and streamflow are used to assess the 420 water requirements for terrestrial river ecosystems. When adhering to EFRs the global water 421 withdrawals are reduced substantially, especially due to groundwater withdrawal limitations. This 422 limitation indicates competition between water allocated for anthropogenic uses and environmental 423 purposes. In addition, groundwater withdrawal reductions upstream lead to increased surface water 424 availability downstream. This interaction results in a trade-off between upstream groundwater 425 withdrawals and downstream surface water withdrawals. Note that VIC-WUR does not include non-426 renewable groundwater withdrawals, while these withdrawals would affect baseflow to a lesser degree.

427 It can be discussed to what extent the EFRs for baseflow are too constricting, because it is based on the 428 relatively stringent EFR for streamflow of Richter et al. (2012) (10 % of the natural streamflow). However, in the absence of any other standards, this baseflow standard remains the best available. 429 430 However, the model setup allows for the evaluation of other standards as well. Note that, even when 431 accounting for EFRs for baseflow on a grid scale, withdrawals can still have local and long-term impacts 432 that are not captured by the model. The timing, location and depth of groundwater withdrawals are also 433 important due to their interactions with the local geohydrology, as discussed by Gleeson and Richter 434 (2018).

The newly developed model will be used for the assessment of trade-offs and synergies between the sectors in the water-food-energy nexus. However, there are some challenges when applying the methods





437 as described in our paper to future water-food-energy nexus assessments. Firstly, an holistic approach 438 is needed to assess trade-offs and synergies in the water-food-energy nexus. This approach should 439 account for competition between resources in an integrated way and should also be captured in 440 consistent scenarios. These scenarios should, for example, define future developments of manufacturing 441 subsectors, hydropower and thermoelectric developments and water efficiencies. Secondly, the energy 442 sector should be expanded upon. Currently only a limited number of thermoelectric power plants are 443 included in the analysis while the rest is incorporated in the industrial water demands. By explicitly 444 accounting for the energy water demand one is able to assess the impact of water scarcity on energy 445 generation and manufacturing separately. Lastly, developments are needed to translate water scarcity 446 into production losses. Currently, only the lack of water is assessed, while the interest is in determining 447 the impact on food and energy production. Therefore, new modules have to be developed that estimate 448 energy generation and food production based on the water availability of the VIC-WUR model.

449 **5** Conclusion

The VIC-WUR model introduced in this paper aimed to provide new opportunities for global water resource assessments. Accordingly, several newly developed anthropogenic impact modules were integrated into the VIC-5 macro-scale hydrological model. The additions presented here comprehensively include anthropogenic and environmental water requirements and expand upon the previous efforts of Haddeland et al. (2006b).

The performance of the modules is in line with reported global water withdrawals and results of other hydrological models. While these additions are sufficient for global water resource assessments, further development is required in order to holistically assess trade-offs and synergies in the water-food-energy nexus. A preliminary assessment of environmental flow requirements already shows competition between water resources allocated for human consumption and the environment, from both ground and surface water sources.





461 6 Code availability

462	All code for the VIC-WUR model is freely available at github.com/wur-wsg/VIC (tag VIC-WUR.2.0.0;
463	DOI 10.5281/zenodo.3399450) under the GNU General Public License, version 2 (GPL-2.0). VIC-
464	WUR documentation can be found at vicwur.readthedocs.io. The original VIC model is freely available
465	at github.com/UW-Hydro/VIC (tag VIC.5.0.1; DOI 10.5281/zenodo.267178) under the GNU General
466	Public License, version 2 (GPL-2.0). VIC documentation can be found at vic.readthedocs.io.
467	Documentation and scripts concerning inputs, configurations and analysis used in this study is freely
468	available at github.com/bramdr/VIC-WUR_support (tag VIC-WUR.2.0.0; DOI
469	10.5281/zenodo.3401411) under the GNU General Public License, version 3 (GPL-3.0).

470 7 Appendix

471 7.1 Appendix A: VIC water and energy balance

In VIC each sub-grid computes the water and energy balance individually (i.e. sub-grid do not exchange 472 473 water or energy between one another). For the water balance, incoming precipitation is partitioned 474 between evapotranspiration, surface and subsurface runoff, and soil water storage. Potential 475 evapotranspiration is based on the Penman-Monteith equation without the canopy resistance 476 (Shuttleworth, 1993). The actual evapotranspiration is calculated by two methods, based on whether the 477 land cover is vegetated or not (bare soil). Evapotranspiration of vegetation is constrained by stomatal, 478 architectural and aerodynamic resistances and is partitioned between canopy evaporation and 479 transpiration based on the intercepted water content of the canopy (Deardorff, 1978; Ducoudre et al., 480 1993). Bare soil evaporation is constrained by the saturated area of the upper soil layer. The saturated 481 area is variable within the grid since (as the model name implies) the infiltration capacity of the soil is 482 assumed heterogeneous (Franchini and Pacciani, 1991). Saturated areas evaporate at the potential evaporation rate while in unsaturated areas evaporation is limited. Surface runoff is produced by 483 484 precipitation over saturated areas. Precipitation over unsaturated areas infiltrates into the upper soil layer 485 and drains through the soil layers based on the gravitational hydraulic conductivity equations of Brooks 486 and Corey (1964). In the first and second layer water is available for transpiration, while the third layer





- 487 is assumed to be below the root zone. From the third layer baseflow is generated based on the non-linear 488 Arno conceptualization (Franchini and Pacciani, 1991). Baseflow increases linearly with soil moisture 489 content when the moisture content is low. At higher soil moisture contents the relation is non-linear, 490 representing subsurface storm-flows.
- 491 For the energy balance, incoming net radiation is partitioned between sensible, latent, and ground heat 492 fluxes and energy storage in the air below the canopy. The energy storage below the canopy is omitted 493 if it is considered negligible (e.g. the canopy surface is open or close to the ground). The latent heat flux 494 is determined by the evapotranspiration as calculated in the water balance. The sensible heat flux is 495 calculated based on the difference between the air and surface temperature and the ground heat flux is 496 calculated based on the difference between the soil and surface temperature. Since the incoming net 497 radiation is also a function of the surface temperature (specifically the outgoing longwave radiation), the surface temperature is solved iteratively. Subsurface ground heat fluxes are calculated assuming an 498 499 exponential temperature profile between the surface and the bottom of the soil column, where the bottom 500 temperature is assumed constant. Later model developments included options for finite difference 501 solutions of the ground temperature profile (Cherkauer and Lettenmaier, 1999), spatial distribution of 502 soil temperatures (Cherkauer and Lettenmaier, 2003), a quasi-2-layer snow-pack snow model 503 (Andreadis et al., 2009), and blowing snow sublimation (Bowling et al., 2004).

504 7.2 Appendix B: EFRs for streamflow and baselow

505 VIC-WUR used the Variable Monthly Flow (VMF) method (Pastor et al., 2014) to limit surface water 506 withdrawals. The VMF method (Pastor et al., 2014) calculates the EFRs for streamflow as a fraction of 507 the natural flow during high (Eq. A.1), intermediate (Eq. A.2) and low (Eq. A.3) flow periods. The 508 presumptive standard Gleeson and Richter (2018) is used to limit groundwater withdrawals. This 509 standard calculates the EFRs for baseflow as 90 % of the natural subsurface runoff through time (Eq. 510 A.4). Here, daily instead of monthly EFRs were used to better capture the monthly flow variability.

511
$$EFR_{s,d} = 0.6 \cdot NF_{s,d}$$
 Eq. (A.1)

512 where
$$NF_{s,d} \leq 0.4 \cdot NF_{s,v}$$

22





513

$$EFR_{s,d} = 0.45 \cdot NF_{s,d}$$
 Eq. (A.2)

 514
 where $0.4 \cdot MF_{s,y} < NF_{s,d} \le 0.8 \cdot NF_{s,y}$
 Eq. (A.3)

 515
 $EFR_{s,d} = 0.3 \cdot NF_{s,d}$
 Eq. (A.3)

 516
 where $NF_{s,d} > 0.8 \cdot NF_{s,y}$
 Eq. (A.4)

 517
 $EFR_{b,d} = 0.9 \cdot NF_{b,d}$
 Eq. (A.4)

518 Where $EFR_{s,d}$ is the daily EFRs for streamflow [m³ s⁻¹], $EFR_{b,d}$ the daily EFRs for baseflow [m³ s⁻¹], 519 $NF_{s,d}$ is the average natural daily streamflow [m³ s⁻¹], and $NF_{s,y}$ is the average natural yearly streamflow 520 [m³ s⁻¹], and $NF_{b,d}$ is the average natural daily baseflow [m³ s⁻¹].

521 EFRs for streamflow and baseflow were based on VIC-WUR naturalized simulations between 1980 and
522 2010. Average natural daily flows were calculated as the multi-year daily average flow over the
523 simulation period, followed by a 30-day moving average smoother.

524 7.3 Appendix C: Dam operation scheme

525 VIC-WUR used a dam operation scheme based on Hanasaki et al. (2006). Target release (i.e. the 526 estimated optimal release) was calculated at the start of the operational year. The operational year starts 527 at the month where the inflow drops below the average annual inflow, and thus the storage should be at 528 its desired maximum. The scheme distinguished between two dam types: (1) dams that did not account 529 for water demands downstream (e.g. hydropower dams or flood control) and (2) dams that did account 530 for water demands downstream (e.g. irrigation dams). The original scheme of Hanasaki et al. (2006) 531 also accounts for EFRs, which were fixed at half the annual mean inflow. Other studies lowered the 532 requirements to a tenth of the mean annual inflow, increasing irrigation availability and preventing 533 excessive releases (Biemans et al., 2011; Voisin et al., 2013). In our study the original dam operation 534 scheme was adapted slightly to account for monthly varying EFRs.

For dams that did not account for demands, the initial release was set at the mean annual inflow corrected
by the variable EFRs (Eq. A.1). For dams that did account for demands, the initial release was increased
during periods of higher water demand. If demands were relatively high compared to the annual inflow,





- 538 the release was corrected by the demand relative to the mean demand (Eq. A.2). If demands were 539 relatively low compared to the annual inflow, release was corrected based on the actual water demand 540 (Eq. A.3).
- 541

542
$$R'_{m} = EFR_{s,m} + (I_{y} - EFR_{s,y})$$
 Eq. (A.1)

543 where $D_y = 0$

544
$$R'_m = EFR_{s,m} + (I_y - EFR_{s,y}) * \frac{D_m}{D_y}$$
 Eq. (A.2)

545 where
$$D_y > 0$$
 and $D_y > (I_y - EFR_{s,y})$

546
$$R'_m = EFR_{s,m} + (I_y - EFR_{s,y}) - D_y + D_m$$
 Eq. (A.3)

547 where
$$D_y > 0$$
 and $D_y \le (I_y - EFR_{s,y})$

548 Where R'_m is the initial monthly target release $[m^3 \ s^{-1}]$, $EFR_{s,m}$ is the average monthly EFR for 549 streamflow demand $[m^3 \ s^{-1}]$, I_y is the average yearly inflow $[m^3 \ s^{-1}]$, $EFR_{s,y}$ is the average yearly EFR 550 for streamflow $[m^3 \ s^{-1}]$, D_m is the average monthly water demand $[m^3 \ s^{-1}]$ and D_y is the average yearly 551 water demand $[m^3 \ s^{-1}]$.

As in Hanasaki et al. (2006), the initial target release was adjusted based on storage and capacity. Target release was adjusted to compensate differences between the current storage and the desired maximum storage (Eq. A.4). Target release was additionally adjusted if the storage capacity is relatively low compared to the annual inflow, and unable to store large portions of the inflow for later release (Eq. A.5).

557
$$R_m = k \cdot R'_m$$
 Eq. (A.4)

558 where $c \ge 0.5$

559
$$R_m = \left(\frac{c}{0.5}\right)^2 \cdot k \cdot R'_m + \left\{1 - \left(\frac{c}{0.5}\right)^2\right\} \cdot I_m$$
 Eq. (A.5)

560 where
$$0 \le c \le 0.5$$





- 561 Where I_y is the average monthly inflow [m³ s⁻¹], *c* the capacity parameter [-] calculated as the storage 562 capacity divided by the mean annual inflow and *k* the storage parameter [-] calculated as current storage 563 divided by the desired maximum storage. The desired maximum storage was set at 85% of the storage 564 capacity as recommended by Hanasaki et al. (2006).
- Water inflow, demand and EFRs were estimated based on the average of the past five years. Water demands were based on the water demands of downstream cells. Only a fraction of water demands were taken into account, based on the fraction of upstream area the dam controlled. For example: if a dam controlled 70% of the upstream area of a downstream cell, than 70% of its demands were taken into account. Fractions smaller than 25% were ignored.
- 570 The original dam operation scheme of Hanasaki et al. (2006) was shown to produce excessively high 571 discharge events due to overflow releases (Masaki et al., 2018). These overflow releases occurred due 572 to a mismatch between the expected and actual inflow. In our study, dam release was increased during 573 high-storage events to prevent overflow and accompanying high discharge events. If dam storage was 574 above the desired maximum storage, target dam release was increased to negate the difference (Eq. A.6). 575 If dam storage was below the desired minimum storage, release is decreased (Eq. A.7). Dam release was 576 adjusted exponentially based on the relative storage difference: small storage differences were only 577 corrected slightly, but if the dam was close to overflowing or emptying, the difference was corrected 578 strongly.

579
$$R_a = R_m + \frac{(S - C\alpha)}{\gamma} \cdot \left(\frac{\frac{S}{c} - \alpha}{1 - \alpha}\right)^b$$
 Eq. (A.6)

580 where
$$S > C\alpha$$

581
$$R_a = R_m + \frac{(S - C(1 - \alpha))}{\gamma} \cdot \left(\frac{(1 - \alpha) - \frac{S}{c}}{1 - \alpha}\right)^b$$
Eq. (A.7)

582 where
$$S < C(1 - \alpha)$$

583 Where R_a is the actual dam release $[m^3 s^{-1}]$, S the dam storage capacity $[m^3]$, α the fraction of the capacity 584 that is the desired maximum [-], β the exponent determining the correction increase [-] and γ the





- 585 parameter determining the period when the release is corrected [s⁻¹]. In testing the exponent and period
- 586 were tuned to 0.6 and 5 days respectively.
- 587 7.4 Appendix D: Water demand

588 7.4.1 Fitting and validation data

Data on irrigation, domestic and industrial water withdrawals were based on the AQUASTAT database
(FAO, 2016), EUROSTAT database (EC, 2019) and United Nations World Water Development Report
(Connor, 2015). Data on GDP per capita and GVA was abstracted from the Maddison Project Database
2018 (Bolt et al., 2018), Penn World Table 9.0 (Feenstra et al., 2015) and World Bank Development
Indicators (World bank, 2010).

594 Available data for domestic an industrial withdrawals were divided into a dataset used for parameter 595 fitting and a dataset used for validation. Domestic water demands were estimated for each United Nations sub-region, and thus data was divided per sub-region to ensure a good global coverage of data. 596 597 In the same manner industrial water demand were divided per country. In case there is only a single data 598 point, the data was added to both the fitting and validation data. To assess uncertainty introduced by 599 dividing the dataset, a sensitivity analysis was performed. The dataset was divided into partial datasets 600 100 times at random. Each partial dataset was used to generate and map domestic and industrial water demands. Domestic water demands based on the partial datasets had a standard deviation of around 9% 601 602 compared to water demands based on the total dataset (Figure A.1a). Industrial water demands based on the partial datasets had a standard deviation of around 3% compared to water demands based on the total 603 604 dataset (Figure A.1b).







Figure A1: Sensitivity analysis of dividing the available data into a fitting and validation dataset for the domestic (a) and industrial (b) water demand estimations. Uncertainty bands are based on the standard deviation of 100 divisions.
The solid line is the water demand estimation based on the total data. The striped line is the water demand estimation used in this study.

610 7.4.2 Irrigation sector

611 Conventional irrigation demands were calculated when soil moisture contents drop below the critical 612 threshold where evapotranspiration will be limited. Demands were set to fill the soil up to field capacity 613 (Eq. A.12). Paddy irrigation demands were set to always keep the soil moisture content of the upper soil 614 layer saturated (Eq. A.13), similar to Hanasaki et al. (2008) and Wada et al. (2014). For paddy irrigation, 615 the saturated hydraulic conductivity of the upper soil layer was reduced by its cubed root to simulate 616 puddling practices, as recommended by the CROPWAT model (Smith, 1996). Total irrigation demands 617 were adjusted by the irrigation efficiency (Eq. A.14). Paddy irrigation used an irrigation efficiency of 1 618 since the water losses were already incorporated in the water demand calculation.

$$\begin{array}{ll} 619 & ID_{conventional}' = FC - (W_1 + W_2) & \text{Eq. (A.12)} \\ 620 & where W_1 + W_2 < W_{cr,1} + W_{cr,2} \\ 621 & ID_{paddy}' = W_{max,1} - W_1 & \text{Eq. (A.13)} \end{array}$$

$$622 \qquad where W_1 < W_{max,1}$$

623
$$ID = ID' * IE$$
 Eq. (A.14)





Where *ID* 'conventional is the conventional crop irrigation demand [mm], *ID* 'paddy is the paddy crop irrigation demand [mm], *ID* is the total irrigation demand [mm], W_1 and W_2 are the soil moisture contents of the first and second soil layer respectively [mm], W_{cr} is the critical soil moisture content [mm], *FC* is the field capacity [mm], W_{max} the maximum soil moisture content [mm], and *IE* is the irrigation efficiency [mm mm⁻¹]. The field capacity was tuned to tuned to (Wcr₁ + Wcr₂) / 0.8. Note that our study used a lower field capacity compared to the (Haddeland et al., 2006b) model setup, since this provided a better fit.

631 7.4.3 Domestic sector

Domestic water demands were represented by using a sigmoid curve for the calculation of structural
domestic water demands (Eq.A.15) and a efficiency rate for the calculation of water-use efficiency
increases (Eq. A.16). These equations differ slightly from Alcamo et al. (2003) since our study used the
base 10 logarithms of GDP and water withdrawals per capita as they provided a better fit.

636
$$DSW_y = DSW_{min} + (DSW_{max} - DSW_{min}) * \frac{1}{1 + e^{-f(GDP_y - o)}}$$
 Eq. (A.15)

637
$$DW_y = 10^{DSW_y} \cdot TE^{y-y_{base}}$$
 Eq. (A.16)

Where *DSW* is the yearly structural domestic withdrawal [log10 m³ cap⁻¹], *DW* the yearly domestic withdrawal [m³ cap⁻¹], *DW*_{min} the minimum structural domestic withdrawal [log10 m³ cap⁻¹], *DW*_{max} the maximum structural domestic withdrawal (without technological improvement) [log10 m³ cap⁻¹], *GDP* the yearly gross domestic product [log10 USD_{equivalent} cap⁻¹], *f* [-] and *o* [log10 USD_{equivalent}] the parameters that determine the range and steepness of the sigmoid curve, *y* the year index, *TE* the technological efficiency rate [-], and *y*_{base} the base year (taken to be 1980).

 DW_{min} was set at 7.5 l cap⁻¹ d⁻¹ based on the World Health Organisation standard (Reed and Reed, 2013), DW_{max} was estimated at around 450 l cap⁻¹ y⁻¹ based on a global curve fit, and *TE* was set at 0.995, 0.99, and 0.98 for developing, transition and developed countries respectively (United Nations development status classification) based on Flörke et al. (2013). Curve parameters f and o were estimated for the 23 United Nations sub-regions based on the GDP per capita and domestic water withdrawal data. In case





- 649 insufficient data was available to calculate parameters values, regional (4 sub-regions) or global (5 sub-
- 650 regions) parameter estimates were used.

651 7.4.4 Industrial sector

- 652 Industrial water demands were represented by using a linear formula for the calculation of structural
- 653 industrial water demands (Eq. A.17) and a efficiency rate for the calculation of water-use efficiency
- 654 increases (Eq. A.18).

$$655 \quad ISW_y = ISW_{int} \cdot GVA_y \qquad \qquad \text{Eq. (A.17)}$$

$$656 \quad IW_y = ISW_y \cdot TE^{y-y_{base}} \qquad \qquad Eq. (A.18)$$

- Where *ISW* is the yearly structural industrial withdrawal $[m^3]$, IW_{int} the country specific industrial water intensity $[m USD_{equivalent}^{-1}]$, IW the yearly industrial withdrawal $[m^3]$, GVA the yearly gross value added by industry $[USD_{equivalent}]$, y the year index, y_{base} the base year (taken to be the year when the industrial water intensity is determined), and *TE* the technological efficiency rate [-].
- TE was set at 0.976 and 1 for OECD and non-OECD countries respectively before the year 1980, 0.976 between the years 1980 and 2000 and 0.99 after the year 2000 based on Flörke et al. (2013). Industrial water intensities were estimated for the 246 United Nations countries based on the GVA and industrial water withdrawal data. In case insufficient data was available to calculate the industrial water intensities, either sub-regional (76 countries), regional (17 countries) or global (6 countries) intensities estimates were used.

667 7.4.5 Energy sector

For each thermoelectric power plant the water intensity was combined with their generation to calculate
the water demands (Eq. A.19). Since there was no observed data about the actual annual generation,
each plant was assumed to be running at its installed generation capacity throughout the year, similar to
van Vliet et al. (2016).





- 673 Where *EW* is the yearly energy withdrawal $[m^3]$, *EW*_{int} the energy water intensity $[m^3 \text{ MWh}^{-1}]$, *G* the
- 674 yearly generation for each plant [MWh], and *y* the year index.
- The energy water demands were subtracted from the industrial water demands at the location of each
- power plant. In cases where the grid cell industrial water demand was less than the energy water demand,
- 677 national industrial water demands were lowered. In cases where even the national industrial water
- 678 demands were less than the national energy water demand (5 countries), the energy water demands were
- 679 lowered instead. This could be the case in countries where power plants do not operate at their installed
- 680 capacity, as globally around 45% of the installed capacity is actually generated (based on data of van
- 681 Vliet et al. (2016)). Energy demands were lowered until 10% of the national industrial water demand
- remains, to ensure some spatial coverage of industrial and energy water demands.

683 7.4.6 Livestock sector

684 Livestock water demands were estimated by combining the livestock population with the water685 requirements for each livestock variety (Eq. A.20).

- 687 Where LW is the yearly livestock withdrawal $[m^3]$, LW_{int} the livestock water intensity $[m^3 \text{ livestock}^{-1}]$,
- 688 L the livestock number for each variety [livestock].

689 8 Author contribution

Bram Droppers and Wietse H.P. Franssen developed and tested the model additions introduced in VICWUR. Bram Droppers generated and analysed the results. Michelle T.H. van Vliet, Bart Nijssen and
Fulco Ludwig provided overall oversight and guidance. Bram Droppers prepared the manuscript with
contributions from all co-authors.

694 9 Competing interests

695 The authors declare that they have no conflict of interest.





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