

## **Reply to reviewer's comments**

Typographical convention:

Black -> reviewer's comment

Blue -> author's comment

Red -> modification to the manuscript

Underlined -> text added to the original version

~~Strikethrough~~ -> text removed from the original version

The line numbers refer to the original version of the manuscript.

### **Reviewer 1 (RC1)**

Realistic representation of urban vegetation is an important step to improving model capabilities and performance. In this paper, a vegetated urban canopy model incorporating street trees (BEP- Tree) is coupled with a mesoscale model (COSMO). The authors briefly describe the relevant parts of the models and their combination, before evaluating the combined model against a range of observations. The work is well-presented, with a clear structure and, mostly, an appropriate level of detail (although in some places more discussion would be beneficial). The inclusion of street trees within the model is shown to make generally small improvements to the meteorology but substantial improvements to the sensible and latent heat fluxes. The coupled model will be useful for applications concerning urban greening as well as for more general studies which will benefit from better representation of vegetation processes.

The paper is very well-presented overall. There are a few consistency issues which need to be addressed and a little more explanation would be beneficial in parts. However, I recommend this publication following minor revisions.

We are happy that the reviewer appreciated the paper. We included further discussion and fixed the inconsistencies.

#### **Main points**

- P8 L15: The lack of interaction with soil moisture seems to be the major weakness in this study. More discussion would be beneficial. Why was the decision made not to implement this interaction? Can the authors comment on the effect of this decision? How realistic is it to assume transpiration in urban areas is never limited by soil moisture (I would think not very realistic)?

We added a discussion on the lack of interaction with soil moisture.

P8 L15 "Although moisture exchange between street trees and the atmosphere is implemented, street trees do not interact with soil moisture content as represented by COSMO's land surface scheme. In other words, a mechanistic interaction between soil moisture and the transpiration of street trees is

not included, assuming that the transpiration is never limited by soil water availability. A careful representation of soil moisture in the urban tiles would have required the development of a new urban hydrology scheme, which was beyond the scope of the study. The missing interaction with soil moisture may reduce the model ability to represent variations in transpiration during periods with large changes in soil moisture (Konarska et al., 2016; Asawa et al., 2017). Nevertheless, street trees are less sensitive to variations in soil moisture than short vegetation, thanks to their deeper root system (Chen et al., 2011; Asawa et al., 2017)."

Additionally, the section "Future studies" is extended accordingly.

P26, L18 "In order to improve the representation of the transpiration from street trees and therefore the modelling of the associated latent heat fluxes, a mechanistic stomata model (e.g. Damour et al., 2010) should be implemented. Furthermore, to properly represent soil water availability during extended drought periods, an urban hydrology module should be developed and coupled to COSMO-BEP-Tree (e.g. Järvi et al., 2011, Yang et al., 2015, Stavropoulos-Laffaille et al., 2018)."

#### Additional references:

Chen, Lixin, et al. "Biophysical control of whole tree transpiration under an urban environment in Northern China." *Journal of Hydrology* 402.3-4 (2011): 388-400.

<https://doi.org/10.1016/j.jhydrol.2011.03.034>

Asawa, Takashi, Tomoki Kiyono, and Akira Hoyano. "Continuous measurement of whole-tree water balance for studying urban tree transpiration." *Hydrological Processes* 31.17 (2017): 3056-3068.

<https://doi.org/10.1002/hyp.11244>

Järvi, Leena, C. S. B. Grimmond, and Andreas Christen. "The surface urban energy and water balance scheme (SUEWS): Evaluation in Los Angeles and Vancouver." *Journal of hydrology* 411.3-4 (2011): 219-237. <https://doi.org/10.1016/j.jhydrol.2011.10.001>

Yang, Jiachuan, et al. "Enhancing hydrologic modelling in the coupled weather research and forecasting–urban modelling system." *Boundary-Layer Meteorology* 155.1 (2015): 87-109.

<https://doi.org/10.1007/s10546-014-9991-6>

Stavropoulos-Laffaille, Xenia, et al. "Improvements to the hydrological processes of the Town Energy Balance model (TEB-Veg, SURFEX v7. 3) for urban modelling and impact assessment." *Geoscientific Model Development* 11.10 (2018): 4175-4194. <https://doi.org/10.5194/gmd-11-4175-2018>

Konarska, Janina et al. "Transpiration of urban trees and its cooling effect in a high latitude city." *International journal of biometeorology* 60 (2016): 159–172. <https://doi.org/10.1007/s00484-015-1014-x>

- P10 L6: It is not clear how  $T_g$  relates to  $T_{nat}$  or  $T_s$  relates to  $T_{urb}$

We apologize that there was a typo in the second part of the equation.  $f_{nat}$  should always go with  $T_g$  and  $f_{urb}$  with  $T_s$ . Additionally, the definition of  $T_s$  was not consistent with its use later in the manuscript (where it is referred to as  $T_{str}$ ).

We corrected the equation and the following explanation.

$$P10, L6 \quad "T_{2m} = (f_{nat} T_g + f_{urb} T_{str}) + r (T_1 - (f_{nat} T_g + f_{urb} T_{str})),$$

where  $T_g$  is the surface temperature of the natural tile,  $T_{str}$  is the street surface temperature ~~of the urban tile, ...~~

- P11 L5: Provide some justification for the statement about Basel's 'abundant green infrastructures' as the statistics and Fig 3b do not suggest a particularly vegetated city.

This statement was based on information provided by the city offices. We reviewed the scientific literature but did not find any studies that substantiate our statement. Therefore, we rephrased the sentence and removed the statement about 'abundant green infrastructures'.

In order to allow a comparison with other studies, we calculated the average canopy cover from street trees and added an according sentence to the manuscript. A figure with the variability of the canopy cover from street trees in the different districts is added to the Supplement (Figure R1).

P11, L3: "The inner city (Basel-Stadt) ~~includes is characterised by an abundance of green areas with~~ more than 24'000 urban trees and 275 hectares of public vegetated surfaces distributed over a total area of the city centre of approx. 2'385 hectares (Stadtgaertnerei Kantons Basel-Stadt). The average canopy cover from street trees is about 20% when considering the entire urban area, but it varies from 6 to 31% in the different districts (see Supplement). The city of Basel is ~~not only~~ an interesting target for ~~its abundant green infrastructures, but also for~~ the availability of extensive observational data sets for model evaluation, which had been used in numerous previous urban climate studies (e.g. Rotach et al., 2005; Parlow et al., 2014)."

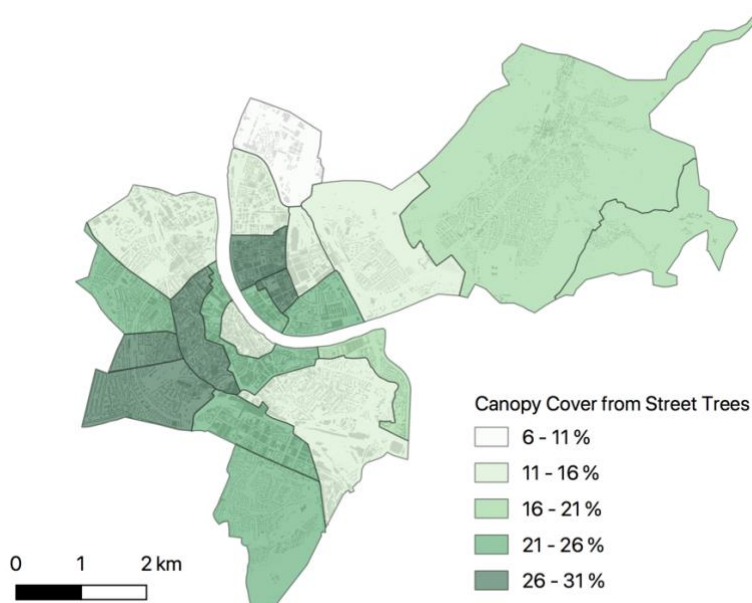


Figure R1 - Canopy cover from street trees in the administrative districts of Basel-Stadt. The building geometries are shown in the background.

- P11 L16: Can the authors add whether any spurious structures occurred (as in other studies, e.g. Salamanca et al. 2012) given the small horizontal grid spacing used is in the ‘grey zone’?  
Salamanca F, Martilli A, Yague C (2012) A numerical study of the urban heat island over Madrid during the DESIREX (2008) field campaign with WRF and an evaluation of simple mitigation strategies, IJC 32: 2372–2386 DOI: 10.1002/joc.3398

Clear spurious structures were not detected in the simulations.

Nevertheless, a sentence regarding the use of the model in the so-called grey-zone is added in the section “future studies”.

P27, L28: “It has to be noted that the model resolution in this study falls within the so-called “grey zone” or “terra incognita” (Wyngaard, 2004) of meso-scale atmospheric modelling, a condition where the horizontal grid spacing and hence the scale of the spatial filter used on the equations of motion are of the same order as the scale of the energy- and flux-containing turbulence. This condition may produce a misrepresentation in the turbulence and wind speed patterns (Efsthathiou et al., 2015). High-resolution urban climate studies reported the formation of horizontal convective rolls which have an unclear counterpart in the real atmosphere (Salamanca et al., 2012; Gutiérrez et al., 2015). Although such issues were not visible in this application, these aspects need to be considered in future studies, for example by exploring the performance of new scale-aware turbulence schemes (e.g. Junshi et al., 2015, Shin et al., 2015) or large-eddy-simulation schemes as recently developed for the COSMO model (Panosetti et al., 2016).”

Additional references:

Efsthathiou, G. A., and Robert J. Beare. "Quantifying and improving sub-grid diffusion in the boundary-layer grey zone." *Quarterly Journal of the Royal Meteorological Society* 141.693 (2015): 3006-3017. <https://doi.org/10.1002/qj.2585>

Gutiérrez, Estatio, et al. "Simulations of a heat-wave event in New York City using a multilayer urban parameterization." *Journal of Applied Meteorology and Climatology* 54.2 (2015): 283-301. <https://doi.org/10.1175/JAMC-D-14-0028.1>

Shin, Hyeyum Hailey, and Song-You Hong. "Representation of the subgrid-scale turbulent transport in convective boundary layers at gray-zone resolutions." *Monthly Weather Review* 143.1 (2015): 250-271. <https://doi.org/10.1175/MWR-D-14-00116.1>

Ito, Junshi, et al. "An extension of the Mellor–Yamada model to the terra incognita zone for dry convective mixed layers in the free convection regime." *Boundary-layer meteorology* 157.1 (2015): 23-43. <https://doi.org/10.1007/s10546-015-0045-5>

Panosetti, Davide, et al. "Idealized large-eddy and convection-resolving simulations of moist convection over mountainous terrain." *Journal of the Atmospheric Sciences* 73.10 (2016): 4021-4041. <https://doi.org/10.1175/JAS-D-15-0341.1>



Salamanca, Francisco, Alberto Martilli, and Carlos Yagüe. "A numerical study of the Urban Heat Island over Madrid during the DESIREX (2008) campaign with WRF and an evaluation of simple mitigation strategies." *International Journal of Climatology* 32.15 (2012): 2372-2386.  
<https://doi.org/10.1002/joc.3398>

Wyngaard, John C. "Toward numerical modeling in the "Terra Incognita"." *Journal of the atmospheric sciences* 61.14 (2004): 1816-1826. [https://doi.org/10.1175/1520-0469\(2004\)061<1816:TNMITT>2.0.CO;2](https://doi.org/10.1175/1520-0469(2004)061<1816:TNMITT>2.0.CO;2)

- P12 Fig 3: Station WEIL does not seem to be used anywhere in the manuscript – delete from the map

The station WEIL is removed from Figure 3.

- P13 L3-4: Two sentences providing a general description of the imperviousness dataset and the buildings dataset would be helpful (i.e. are these fractions of impervious surface cover or degree of imperviousness; is it building height or building material or building use; at what resolution?)

We added two sentences describing the imperviousness dataset and the building dataset.

P13, L5: "The urban canopy parameters were derived with UCPgenerator from the following input datasets: (a) imperviousness from EEA (2015), (b) 3D buildings from Federal Office of Topography (2007) and (c) lidar-based tree canopy height (see data availability section). The imperviousness dataset represents the percentage of soil sealing (including building area) and has a resolution of 20 m. The building dataset is a vector dataset representing the building geometries (accuracy 3-8 m), including heights but without information on the building materials or use."

- P15 Table 3: Latent heat flux seems to be denoted  $Q_L$  here and  $Q_E$  later in the manuscript – make consistent in all text, tables and figures. Fluxes also appear in the row for BLER – presumably this is incorrect?

Thank you for pointing out this. We removed the inconsistency in latent heat flux notation throughout the manuscript. We removed the fluxes from Table 3.

- P15 L21-22: 'net' → 'four-component' as the CNR4 measures all four radiation components individually. Also, the radiation measurements are mentioned here but not used in the evaluation (e.g. in Fig 5). Why not? Can the authors comment on the results for the radiation components for STD and LAO runs?

Thank you for pointing this out. We corrected the description of the CNR4 measurement.

P15, L21-22: "radiation fluxes are measured with a ~~net~~ four-component radiometer"

The comparison with radiation is only shown in the Supplement, since we expected only a minor impact

from street trees.

We added the following explanation regarding the evaluation in terms of radiation fluxes:

P20, L12: “For completeness, the comparison between modelled and observed radiation fluxes is shown in the Supplement. The model simulations agree reasonably well with the observed radiation fluxes, although the upward short-wave radiation is underestimated. This indicates that the albedo is underestimated, which can be related to the choice of material properties (Table 2). The upward long-wave radiation is slightly overestimated, which can therefore be related to the underestimation of the albedo and, consequently, to an overestimation of surface temperatures. Overall, the representation of street trees had a very minor impact on the radiation fluxes.”

- P16 L4: Aren't there only five surface stations (depending on whether the WEIL data should be included or not)? Can the authors clarify here why no results from WEIL have been included?

Thank you for pointing out this inconsistency.

There are indeed only five sites since the site of WEIL has not been used.

P16, L4: ~~“Six~~ Five surface stations are used ...”

- P17 L5: Here the emissivity values for urban surfaces are given as 0.95-0.97 but in Table 2 much lower values of 0.85, 0.9 and 0.95 are given, which seems to introduce an inconsistency. How would an emissivity of 0.85 affect the uncertainty estimation in L7?

Thank you for pointing out this inconsistency.

We recalculated the LST with the values of emissivity as used in the model from Loridan and Grimmond (2012) and compared with those recommended by Sobrino et al., (2008) and Baldrige et al., (2009). We added an additional figure in the Supplement (Figure R1) that highlights the impact of this choice on the LST and edited the text with a comment on the uncertainty introduced by this aspect.

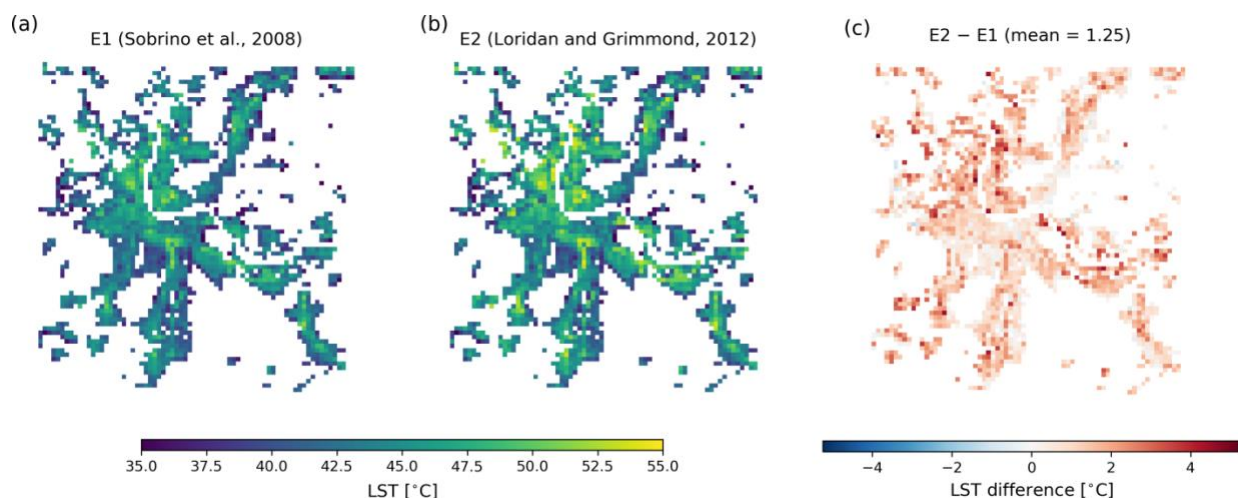


Figure R2 – Impact of the choice of emissivity values for urban surfaces on the land surface temperature (LST). (a,b) show the observed LST with emissivity values from Sobrino et al., (2008) and Loridan and Grimmond (2012), respectively; (c) shows the difference in LST between the two approaches.

“A modified vegetation-threshold approach by Sobrino et al., (2008) was used to estimate the surface emissivity. As the original approach does not account for built-up areas, a lookup table for dark and bright urban surfaces with a range of 0.95 to 0.97, based on emissivity values for typical urban fabrics (Baldrige et al., 2009), was applied (Mitraka et al., 2012). Note that these values differ from those used in the model (Table 2). The uncertainty in the values of emissivity produces average changes in the LST of about 1.25 K (Supplement). On top of this, the LST is affected by uncertainties due to atmospheric correction, sensor and off-nadir view (Voogt and Oke, 2003; Chen et al., 2017) Using the above-described methods, the uncertainty in surface emissivity and atmospheric effects results in a total error of approximately 0.4 to 1.1 K (Jiménez-Muñoz and Sobrino, 2006).”

- P17 L15-17: Define  $f_{dir}$  here as it does not seem to be defined in the main text

We added a definition of  $f_{dir}$ .

P17, L15-16: “...,  $f_{dir}$  is the fraction of canyons with direction  $d$ , ...”

- P17 Fig 5 caption: The letters referring to the various panels needs correcting here and in the main text.

P18, Fig5 caption: “Comparison between observations (Obs) and model simulations (STD, LA0) of air temperature (a,b), specific humidity (b,c,d), wind speed (e,f), sensible heat flux (g,h) and latent heat flux (i,j) at the site BKLI (38 m tall flux tower) during the selected simulation period (discarding the first 5 days as spin-up). The left (a,c,e,g,i) and right (b,d,f,h,j) columns show the 3-h moving average and the diurnal cycle, respectively. ...”

P18, L1-2: “The model simulates the evolution of air temperature very well during the evaluation period (Figure 5a-b) ...”

P19, L1: “The evolution of specific humidity  $q$  throughout the simulation is generally well reproduced (Figure 5a-b,c-d).”

P19, L9: “The model simulates well the evolution of wind speed ( $u$ ) during the evaluation period (Figure 5e-d e-f), ...”

P19, L16: “The sensible heat flux ( $Q_H$ ) follows very well the observations during the evaluation period (Figure 5e-g-h), ...”

P20, L4: “The evolution of latent heat flux ( $Q_E$ ) is generally well reproduced by the model (Figure 5g-h-i-j), ...”

- P17 Fig 5 caption and in other captions: Personally I find it quite confusing to keep reading about

the 5 days that were discarded for spin-up. Why not define a simulation period and a study period/analysis period (27 June-9 July) in Section 3.1, and then refer to this study period in the figure and table captions? Repetitions between captions could also be reduced.

Following this suggestion, we applied the following changes to the manuscript:

P11, L9: “In this study, COSMO-BEP-Tree was run for the period 22 June 2015 - 9 July 2015, with the first 5 days discarded as spin-up period. The analyses were performed on the period 27 June 2015 - 9 July 2015, called study period hereinafter.”

P18, Figure 5 caption: “... during the ~~selected simulation study period (discarding the first 5 days as spin-up)~~. ...”

P18, Table 4 caption: “... during the study period ~~22 June 2015 - 9 July 2015 (discarding the first 5 days as spin-up)~~. ...”

P21, Figure 6 caption: “...during the ~~selected~~ study period ~~(discarding the first 5 days as spin-up)~~. ...”

P22, Table 5 caption: “during the study period ~~22 June 2015 - 9 July 2015 (discarding the first 5 days as spin-up)~~. ...”

P24, Table 6 caption: “during the study period ~~22 June 2015 - 9 July 2015 (discarding the first 5 days as spin-up)~~. ...”

- P19 Table 4 caption: Some inconsistency with the symbol for temperature that needs to be fixed (this was  $T$  previously, and is now  $T_{air}$ )

Table 4 is updated with  $T$  instead of  $T_{air}$ .

- P19 L6: Suggest changing ‘a small increase’ to ‘a very small increase’ and presumably ‘relative humidity’ should be ‘specific humidity’ seeing as the units given are  $\text{g kg}^{-1}$

We agree with the reviewer and changed the text accordingly.

- P19 L10: The wind speed overestimation occurs often: suggest changing ‘a few days’ to ‘most days’

We agree with the reviewer and changed the text accordingly.

- P19 L9-12: The modelled wind speed is large, often exceeding  $5 \text{ m s}^{-1}$ . This seems quite large for a thermally driven circulation between city and surroundings. Perhaps the prescribed roughness lengths could also be a relevant factor. Did the authors consider this and could they comment on the impact of uncertainty in the prescribed roughness lengths on the evaluation results?

First of all, we add the prescribed values of roughness length to Table 2. The values correspond to the

default options of the urban canopy model BEP, as used in previous studies (Chen et al., 2011; Salamanca et al., 2011; Hammerberg et al., 2018).

|  | Roof        | Wall     | Road        |
|--|-------------|----------|-------------|
| Albedo (-)   | 0.1         | 0.1      | 0.15        |
| Emissivity (-)   | 0.85        | 0.9      | 0.95        |
| Heat Capacity ( $\text{J m}^{-3} \text{K}^{-1} \times 106$ ) | 1.5         | 1.4      | 1.5         |
| Heat Conductivity ( $\text{W m}^{-1} \text{K}^{-1}$ )        | 0.8         | 1.0      | 0.8         |
| Total thickness (m)  | 0.5         | 0.3      | 1.0         |
| <u>Roughness length (m)</u>                                  | <u>0.01</u> | <u>-</u> | <u>0.01</u> |

Table 2. Material properties used in the urban canopy model.

We agree with the reviewer that wind speed may be sensitive to values of roughness length at individual surfaces. However, we haven't performed any sensitivity studies on the impact of the prescribed values of roughness length. Moreover, we are also not aware of any studies in the literature where this aspect has been investigated in the context of mesoscale modelling with multi-layer urban canopy model.

Nevertheless, a similar overestimation is also observed at the rural site (Figure 1 in the Supplement). In our opinion, this points out to an overestimation of regional scale wind speed rather than local effects in the urban area.

P19, L10: "...~~The origin of the~~ overestimation was not limited to the urban area but was also found at the rural site (Supplement). Although its origin is unclear, possible causes may be a misrepresentation of the synoptic regional scale wind from the driving boundary conditions and an overestimation of the thermally driven flow induced by the surface temperature gradient between the city and the surroundings."

- P20 L21: 'may be related to exposure of the sensor to local influences' – please explain

We extended the explanation regarding the sensor exposure, as already done when describing the sites (P16, L8).

P19, L10: "This difference may be related to the local exposure of the BLER sensor, located in the vicinity of a (warm) south facing wall."

- P21 Fig 6: Why are no results for BLER plotted here? This would help the reader to interpret Fig 7.

We included the daily profile of air temperature at BLER in Figure 6.

- P21 L4-7: A little deeper discussion would be helpful here. The performance of the model seems to go in the wrong direction in terms of comparison against the observations but this is because of the dominating effect of the temperature bias. However, the newly-incorporated processes seem to cause a change in the expected direction. This cautions against tuning models to match

observations

We included a discussion in order to make it clearer that other model biases were masking the benefit of representing street trees for UHI estimation.

P21 L4-7 “The representation of street trees affects the simulated UHI intensity by producing a slight reduction in UHI intensity from late morning to early evening. The impact of street trees is more evident at sites with higher urban fraction and density of street trees. An improvement in model performance with street trees was, however, not evident in terms of UHI representation. Considering that street trees produce changes in the expected direction, this suggests that probably other model biases are masking the benefit of including street trees. The representation of street trees did not improve the model performance in terms of UHI representation, given that the warm bias at the rural site dominates.”

- P22 L10-12: This two-sentence paragraph should be extended to include more discussion (with references from the literature) about the uncertainties of satellite-derived LST, particularly in urban areas. In Fig 8, do you trust the spatial patterns seen in the satellite observations? Crawford et al- (2018) discuss some of the biases affecting LST measurements in urban areas and how these vary spatially.

Crawford B, Grimmond CSB, Gabey A, Marconcini M, Ward HC, Kent CW (2018) Variability of urban surface temperatures and implications for aerodynamic energy exchange in unstable conditions, QJRMS 144: 1719–1741 DOI: 10.1002/qj.3325

We thank the reviewer for the suggestion and extended the paragraph with more discussion about uncertainties in satellite-derived LST in urban areas.

P22 L10-12 “The explicit representation of street trees did not improve the agreement with satellite observations of land surface temperature. However, this observation should be taken with care given, that satellite-derived LST in urban areas are known to have a large degree of uncertainty due to potential errors in the surface emissivity, sensor and satellite view bias (Voogt and Oke, 2003; Chen et al., 2017). Slight off-nadir view angles can introduce a bias over areas with tall buildings while LST variations at scales below 500 m may be unresolved (Crawford et al., 2018). ~~However, this observation should be taken with care given the uncertainty in the satellite observations of the order of 1 K.~~”

#### References:

Voogt, James A., and Tim R. Oke. "Thermal remote sensing of urban climates." *Remote sensing of environment* 86.3 (2003): 370-384. [https://doi.org/10.1016/S0034-4257\(03\)00079-8](https://doi.org/10.1016/S0034-4257(03)00079-8)

Chen, Feng, et al. "Challenges to quantitative applications of Landsat observations for the urban thermal environment." *Journal of Environmental Sciences* 59 (2017): 80-88. <https://doi.org/10.1016/j.jes.2017.02.009>

Crawford, Ben, et al. "Variability of urban surface temperatures and implications for aerodynamic energy exchange in unstable conditions." *Quarterly Journal of the Royal Meteorological Society* 144.715 (2018): 1719-1741. <https://doi.org/10.1002/qj.3325>

- P26 L3-4: Can the authors explain the finding that both increasing and decreasing leaf area density leads to a reduction in night-time temperature?

We thank the reviewer for the suggestion and included an explanation.

P26, L3-4: “A reduction in the leaf area density (LA-) generally produced an opposite response to that of an increase in leaf area density (LA+). An interesting exception is  $T_{air}$  at night, suggesting that different processes with opposing sensitivities to values of LAD are at work. The largest night-time reduction in  $T_{air}$  was achieved with LA-, which was associated with a similar reduction in  $T_{str}$ . A slight reduction in  $T_{air}$ , however, was also found for LA+, which contrasts with an increase in  $T_{str}$  for this scenario. This points towards complex, non-linear interactions between the different effects of LAD on radiation and winds. However, further studies are needed to corroborate this finding.”

#### Minor points and suggestions for improving language and readability

- P2 L7: interactions concerning moisture should also be listed here
- P2 L11: ‘heat interactions’ is probably not the right term here, suggest changing to ‘latent heat’ or ‘evaporation’
- P2 L12: ‘excerpts’ → ‘extracts’
- P3 L4: ‘Several... ..fabrics’ – this sentence is redundant given the previous and following sentences so can be deleted
- P3 L7: ‘regarded’ → ‘concerns’
- P3 L9: ‘regards’ → ‘concerns’
- P3 L11-12 and throughout: I would use the singular, i.e. ‘street tree datasets’ instead of ‘street trees datasets’; ‘street tree characteristics’ instead of ‘street trees characteristics’ (otherwise you probably need an apostrophe)
- P3 L16: Suggest deleting ‘in the presence of street trees’
- P3 L26: Suggest deleting ‘the modelling’ and adding to the end of this sentence ‘by more realistic representation of the surface and its interactions’, or similar
- P4 L4: ‘garden’ → ‘gardens’
- P4 L6-8: There is no mention of Section 4 here which seems odd
- P4 L12, L24: It is not clear that ‘(COSMO)’ is a reference. You could delete it from L12 and perhaps provide a year or give the URL in L24 to avoid confusion.
- P4 L20-22: The items in this list are not strictly all processes, suggest rephrasing.
- P5 Fig1 caption: ‘the Sections’ → ‘Sections’
- P5 L6: Delete ‘top of’
- P6 L4: Delete the second instance of ‘i’ as it seems as though ‘the ray i’ is defined twice
- P6 L25: Seems strange to define the density here before  $S_q$
- P7 L3: ‘latter’ → ‘canyon geometry inputs’ and ‘includes’ → ‘include’
- P7 L16: ‘a measurement site’ → ‘measurements’
- P7 L23-5: This sentence disrupts the flow here. Suggest deleting or making this the first sentence of Section 2.3.
- P7 L26: Would ‘proportions’ be better than ‘configurations’ here? – the model does not account for the arrangement of the surfaces or whether there is one larger urban surface or many smaller urban surfaces with the same total area
- P8 Fig 2 caption: ‘Rose’ → ‘Orange’. Also most of these are quantities not ‘operations’ – suggest rephrasing

- P8 L1: Delete 'street canyon' (as sky is not really part of the street canyon)
- P10 L4: Start this sentence with 'Following Schubert...'
- P10 L13: 'UPC' → 'UCP'
- P10 L14: The reference to Fig 2 would be more helpful in L16
- P11 L1-8: Change the numbers in this section so that commas rather than apostrophes separate the digits (e.g. '830'000' → '830,000')
- P12 Fig 3 caption: State what the dashed line represents in the figure caption
- P13 L1: 'in the centre-south and centre-west' → 'south and west of the centre'
- P14 Fig 4 caption: 'urban tile' is probably be better than 'urban canopy' as the latter could be interpreted as trees above buildings
- P14 Fig4 caption: Change to 'Further details about the variables are given in Table A1'
- P14 Table 2 and in other tables too: Use a consistent number of decimal places, e.g. 0.10 0.10 and 0.15 for albedo; 0.85, 0.90 and 0.95 for emissivity
- P15 L4: Change to 'vehicles and seasonal'
- P15 L7: It would be more useful to explicitly state the period over which the average was calculated
- P15 L13 and elsewhere: Should this be 'MCR-Lab' not 'MRC-Lab'?
- P15 L14: 'building' → 'buildings'
- P15 L17: ', but' → 'and'
- P15 L17: 'estimated to' → 'estimated at' or 'estimated to be'
- P15 L22: You could add 'in general' to make it clear that these considerable uncertainties are not only for the Basel dataset
- P16 L7: 'placed' → 'located'
- P16 L8: 'then' → 'than'
- P16 L8-9: 'in the area' – in which area? The urban canopy layer is extremely spatially variable so there would be considerable temperature variation expected throughout any 'area'
- P16 L19: Delete 'as well'
- P16 L21: Change to 'Given the vegetated environment and distance from the city centre,' to avoid misinterpretation
- P16 L21-22: Change to 'as a rural reference'
- P17 L30: Delete 'will'
- P17 Fig5: Delete the unhelpful minor tick marks on the right-hand plots.
- P19 Table 4 caption: Change '(see Table 1... set-up)' to '(LAO)'
- P19 Table4 caption: You could point out that the units only apply to RMSE and MBE, not  $r^2$
- P19 L11: Change 'and' to 'or'
- P19 L15: Change 'of' to 'by'
- P20 L3: Shouldn't '0.06' be '0.04'?
- P20 L6: Suggest deleting 'slight' as the overestimation is appreciable. Also add that  $r^2$  gets worse here.
- P20 L16: Delete '2-m air temperature'
- P20 L17: Delete 'will'
- P20 L30-31: Avoid this two-sentence paragraph by moving 'The statistical scores...only.' to the end of the previous paragraph and 'The comparison between... Fig. 7.' to the start of the following paragraph.
- P21 L1: 'lower' → 'lesser'
- P21 L14 'outskirt' → 'outskirts'
- P22 L3: Change 'only quite' to 'quite'
- P22 L3: Change to '...STD and LAO are seen, with STD...'
- P23 L2: 'ST' → 'surface temperature'
- P23 L6: Why is 2-m air temperature now denoted  $T_{air}$  and not  $T_{2m}$  as before?
- P24 Table 7: Units of surface temperature are missing
- P26 Fig9 caption: Change 'S' to ' $S_{\varphi}$ ' in line with Eq 15



- P26 L13-4: Join this first sentence with the following paragraph
- P27 L3: Change to 'of the clumping index'
- P27 L6-7: Change to 'the temporal changes of  $\Omega$  with solar zenith angle'
- P27 L24: Change to 'values of the material properties for roof, wall and street elements were used'
- P27 L28: Should read 'WUDAPT'
- P28 L15: Change to: 'generally well, although its magnitude'
- P28 L16: Change 'it' to 'this difference'
- P28 L18-21: Again some mention of uncertainties in LST would be helpful here
- P28 L23: Change to 'found only at some'
- P29 L7: Close bracket missing after the email address
- P29 L9: Change to 'as a standalone'
- P30 L12-3: Define  $A_f$
- P31 L1: 'heigh' → 'height'
- P32 L24 and L32: Check reference formatting

We thank the reviewer for the careful inspection of the manuscript and have adopted it accordingly.

## **Reply to reviewer's comments**

Typographical convention:

Black -> reviewer's comment

Blue -> author's comment

Red -> modification to the manuscript

Underlined -> text added to the original version

~~Strikethrough~~ -> text removed from the original version

The line numbers refer to the original version of the manuscript.

### **Reviewer 2 (RC2)**

This paper describes and evaluates COSMO-BEP-Tree v1.0: a coupled urban climate model with explicit representation of street trees. The authors assess the model during a heatwave event with observations from flux towers, surface stations and satellites. The authors have presented a very strong study which is thorough and convincing. The issue of integrating the effects of street trees in urban meteorological models is important because increasing tree cover is “go-to” response of urban planners who wish to reduce urban heat impacts. As the manuscript shows, interacting processes are complex and outcomes not always obvious. The evaluation is clear and thorough, and the technical achievements will be beneficial for future urban studies.

The lack of open availability of the source code is disappointing for a study presented in GMD. The process for accessing code outlined in the study is not timely, and includes mailing physical documents and/or negotiation. I have therefore not reviewed the code. The authors should reconsider publishing the code openly on a persistent public archive. Detailed comments follow.

We are happy that the reviewer appreciated the paper.

We regret that the source code was not easily accessible. Unfortunately, we have to conform to the CLM-Community licence agreement that only allows sharing the code with registered members. As such, we cannot publish the code on a public archive and we don't have control on the registration process.

However, we agree with the reviewer that the registration procedure must be simplified and made timely. We will make this point to the CLM-community.

Pg 2 Ln 12: “Tree transpiration reduces the surface temperature of the foliage by converting part of the solar radiation to latent heat” not only solar radiation.

We thank the reviewer for pointing out the inaccurate description of the energy exchanged during transpiration.

We edited the sentence as follows:

Pg 2 In 12: ~~“In terms of heat interactions, Through tree transpiration the leaf surface temperature is reduced<sub>s</sub> resulting in the extraction of sensible heat from the air of the foliage by converting the solar radiation to latent heat (Green, 1993).”~~

Pg 2 In 15: “Modelling studies on the cooling potential of street trees have almost exclusively been performed at the scale of individual street canyons to single neighbourhood...Only very few studies investigate the city-wide impact of street trees, mainly due to the limited availability of models able to represent street trees at the scale of the city...” It would be useful here to acknowledge the various mesoscale urban meteorological models which have incorporated vegetation as low-height gardens within the street canyons (e.g. Thatcher and Hurley, 2012; Lemonsu et al, 2012; Wang et al 2013). Although not “street-trees” (as they do not provide shading on walls or reduce canyon sky view), they do shade the ground and interact directly with canyon radiation/turbulent fluxes and momentum budgets, alter the Bowen ratio and reduce canyon surface temperatures. This will also give an opportunity to be clearer about the authors definition of “street trees” when the concept is introduced, as readers may assume models within-canyon low vegetation, or even external vegetation tiles, are doing the same thing. The statement “very few studies investigate the city-wide impact of street trees” could then be more carefully stated, as there have been many studies at city or larger scales which have assessed the impact of urban vegetation, but most have used schemes with low vegetation or used a tiled approach, and hence missed important shading and sky view effect which is the strength of the current study.

\* Thatcher M and Hurley P 2012 Simulating Australian urban climate in a mesoscale atmospheric numerical model Boundary-Layer Meteorol 142 149–75

\* Lemonsu A, Masson V, Shashua-Bar L, Erell E and Pearlmutter D 2012 Inclusion of vegetation in the Town Energy Balance model for modelling urban green areas Geosci. Model Dev. 5 1377–93

\* Wang Z-H, Bou-Zeid E and Smith J A 2013 A coupled energy transport and hydro-logical model for urban canopies evaluated using a wireless sensor network Quarterly Journal of the Royal Meteorological Society 139 1643–57

We thank the reviewer for these excellent suggestions. We re-worded the section in order to acknowledge studies using the tile approach and those explicitly representing low vegetation.

Pg 2 In 15: ~~“Modelling studies on the cooling potential of street trees have almost exclusively been performed at the scale of individual street canyons to a single neighbourhood (e.g. Gromke et al., 2015; Ng et al., 2012). Only very few studies investigated the city-wide impact of street trees, mainly due to the limited availability of models able to represent street trees at the scale of the city or at even larger scales. In fact, the vast majority of meso-scale urban climate models only represent trees vegetation outside the street canyon, neglecting important effects such as the shading effect of trees on the canyon’s surfaces.”~~

Pg 2 ln 15: "The climatic impact of urban vegetation has been investigated in numerous previous studies from the scale of the single street canyon to that of the entire urban region (e.g. Gromke et al., 2015; Ng et al., 2012; De Munck et al., 2018). However, studies on entire urban regions primarily focused on low vegetation, representing low height gardens and green roofs (e.g. Wang et al., 2013; De Munck et al., 2018). Street trees, instead, have generally only been represented in a separate natural tile (e.g. Schubert and Grossman-Clarke, 2013; Li and Norford, 2016). This approach precluded considering any interactions between trees and urban surfaces in a street canyon, such as shading and sheltering effects (Kravenhoff et al. 2014, 2015). Other studies employed somewhat more sophisticated methods but still neglect some of the critical interactions between trees, canyon surface and airflow (e.g. Thatcher and Hurley, 2012)."

New references:

Schubert, Sebastian, and Susanne Grossman-Clarke. "The Influence of green areas and roof albedos on air temperatures during Extreme Heat Events in Berlin, Germany." *Meteorologische Zeitschrift* 22.2 (2013): 131-143. <https://doi.org/10.1127/0941-2948/2013/0393>

De Munck, C., et al. "Evaluating the impacts of greening scenarios on thermal comfort and energy and water consumptions for adapting Paris city to climate change." *Urban Climate* 23 (2018): 260-286. <https://doi.org/10.1016/j.uclim.2017.01.003>

Li, Xian-Xiang, and Leslie K. Norford. "Evaluation of cool roof and vegetations in mitigating urban heat island in a tropical city, Singapore." *Urban Climate* 16 (2016): 59-74. <https://doi.org/10.1016/j.uclim.2015.12.002>

Kravenhoff, E. S., et al. "A multi-layer radiation model for urban neighbourhoods with trees." *Boundary-layer meteorology* 151.1 (2014): 139-178. <https://doi.org/10.1007/s10546-013-9883-1>

Kravenhoff, E. S., et al. "Parametrization of drag and turbulence for urban neighbourhoods with trees." *Boundary-layer meteorology* 156.2 (2015): 157-189. <https://doi.org/10.1007/s10546-015-0028-6>

Thatcher, Marcus, and Peter Hurley. "Simulating Australian urban climate in a mesoscale atmospheric numerical model." *Boundary-layer meteorology* 142.1 (2012): 149-175. <https://doi.org/10.1007/s10546-011-9663-8>

pg 6 ln 4: Why are the intensity terms ( $\Delta V_i$  and  $r_i$ ) dimensionless? Shouldn't they have units?

We apologize that the units were missing. Both terms indeed have units of  $[W m^{-2}]$ .  
The sentence will be changed as follows:

pg 6 ln 4: "where  $\Delta V_i$  is the reduction in intensity of ray  $i$  due to the tree foliage  $[W m^{-2}]$ ,  $r_i$  is the initial intensity of the ray  $i$   $[W m^{-2}]$ , ... "

pg 8 ln 13 "street trees do not interact with soil moisture content as represented by COSMOS's land surface scheme" More explanation as to why this was the case –technical constraints? Or was it

assumed that all soil is constantly irrigated in urban areas? You may wish to mention the drawback of this assumption (i.e. over estimating latent heat) at this time. Otherwise, do the underlying LSM soil and urban scheme exchange any fluxes? How? Figure 1 could be used here to better explain how fluxes of the urban scheme are coupled to the atmosphere and LSM.

This point was also raised by the other reviewer. We added further discussion on the lack of interaction with soil moisture.

Further motivation is provided in the section “Methodology”

pg8 ln 12: “Although moisture exchange between street trees and the atmosphere is implemented, street trees do not interact with soil moisture content as represented by COSMO’s land surface scheme. In other words, a mechanistic interaction between soil moisture and the transpiration of street trees is not included, assuming that the transpiration is never limited by soil water availability. A careful representation of soil moisture in the urban tile would have required a new urban hydrology scheme, which was beyond the scope of the study. The missing interaction with soil moisture may reduce the model ability to represent variations in transpiration during periods with large changes in soil moisture (Konarska et al., 2016; Asawa et al., 2017). Nevertheless, street trees are less sensitive to variations in soil moisture than short vegetation, thanks to their deeper root system (Chen et al., 2011; Asawa et al., 2017).”

#### References:

Chen, Lixin, et al. "Biophysical control of whole tree transpiration under an urban environment in Northern China." *Journal of Hydrology* 402.3-4 (2011): 388-400.

<https://doi.org/10.1016/j.jhydrol.2011.03.034>

Asawa, Takashi, Tomoki Kiyono, and Akira Hoyano. "Continuous measurement of whole-tree water balance for studying urban tree transpiration." *Hydrological Processes* 31.17 (2017): 3056-3068.

<https://doi.org/10.1002/hyp.11244>

Konarska, Janina, et al. "Transpiration of urban trees and its cooling effect in a high latitude city." *International journal of biometeorology* 60.1 (2016): 159-172. <https://doi.org/10.1007/s00484-015-1014-x>

pg 10 ln 6: For the last terms of eq 13, why is  $T_s$  associated with  $f_{nat}$ , and  $T_g$  be with  $f_{urb}$ ? I don't understand whether this temperature averaging approach is valid. I can see the equation is taken from Schubert (2013), but it is not explained there either. For example, take the extreme position where  $f_{nat}=0$ ,  $f_{urb}=1$ , then the equation simplifies to:  $T_{2m} = T_s + r (T_1 - T_g)$  Why is  $T_g$  part of the  $T_{2m}$  temperature if there is no  $f_{nat}$ ? Also, how is the urban surface temperature here defined? Does it, for example, include the roof facets? I was pleased to see the clear definition of surface temperature for the evaluation of the satellite observations (pg 17) which was based on the facets that the satellite sees (roofs streets and natural surfaces rather than walls). But it's not clear that definition is appropriate for 2m temperature out of sight of roofs, and within sight of walls. If the same satellite-based definition is used for the  $T_{2m}$  calculation, authors should note that's likely to lead to underestimation of the cooling benefit of street trees for temperatures within the canyon.

We apologize that the terms were mixed up in the second part of the equation.  $f_{\text{nat}}$  should always go with  $T_g$  and  $f_{\text{urb}}$  with  $T_s$ . Additionally, the definition of  $T_s$  was not consistent with its use later in the manuscript (where it is referred to as  $T_{\text{str}}$ ).

We corrected the formula and the following explanation.

P10, L6  $T_{2m} = (f_{\text{nat}} T_g + f_{\text{urb}} T_{\text{str}}) + r (T_1 - (f_{\text{nat}} T_g + f_{\text{urb}} T_{\text{str}})),$

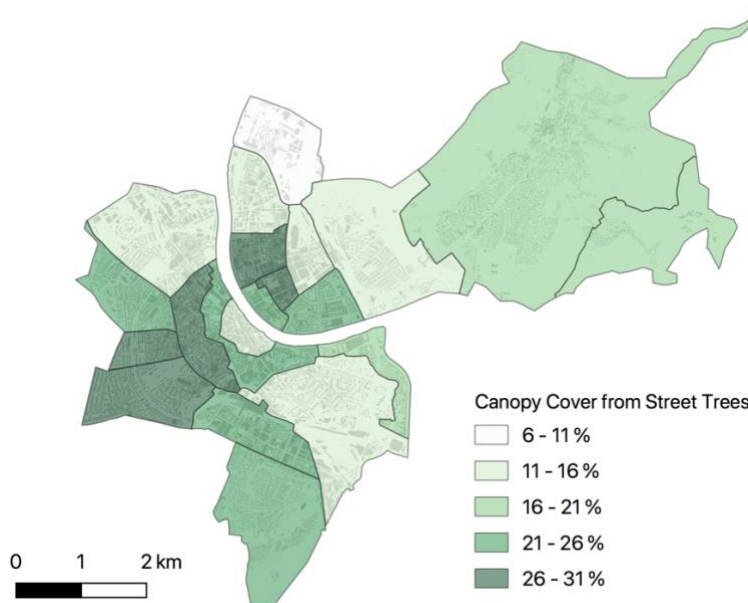
where  $T_g$  is the surface temperature of the natural tile,  $T_{\text{str}}$  is the street surface temperature ~~of the urban tile, ...~~

pg 11 ln 3: It would be useful to include information here about the average fraction of green cover in Basel, for easy comparison with other urban studies.

We thank the reviewer for the suggestion. However, we think that the average fraction of green cover (as normally defined in urban climate studies) is not the best measure, since it includes all urban vegetation (e.g. parks), not only street trees, which is the focus of the present study. Instead, we calculated the average canopy cover from street trees only.

Additionally, we included an additional figure that shows the variability of the canopy cover from street trees in the different districts (Figure R1). The figure is added to the Supplement.

pg 11 ln 5: “The average canopy cover from street trees is about 20% when considering the entire urban area, but it varies from 6 to 31% in the different districts (see Supplement).”



**Figure R1 - Canopy cover from street trees in the administrative districts of Basel-Stadt. The building geometries are shown in the background.**

p 11 ln 23: There is discussion of atmospheric initial conditions and on soil properties but it is not clear how soil moisture was initialised. Soil moisture has significant impact on the intensity of

heatwaves (e.g. see Wang et al 2019), and soil moisture has memory much longer than the 5-day spin up, so this variable needs to be carefully initialised.

\* Wang P, Zhang Q, Yang Y and Tang J 2019 The Sensitivity to Initial Soil Moisture for Three Severe Cases of Heat Waves Over Eastern China Front. Environ. Sci. 7

Soil moisture is indeed a critical quantity in these simulations. Soil moisture (in the non-urban tiles) is computed prognostically for 7 soil layers by the land-surface module TERRA of the COSMO model. Initial soil conditions for our simulations were taken from the operational COSMO-2 analyses of MeteoSwiss. Since the COSMO-2 model (covering the Alpine domain) is nested into the COSMO-7 model (covering Europe), which itself is nested into ECMWF's IFS model, the initial soil moisture field ultimately goes back to the soil moisture analyses of ECMWF's operational IFS model.

We edited the sentence regarding initial and boundary conditions by specifically mentioning that they include soil moisture.

P 11 In 21: "Soil moisture, a key variable for accurate representation of heatwaves (Wang et al., 2019), is also initialised with the COSMO-2 analyses, which trace back to the soil moisture analyses of ECMWF's operational IFS model (De Rosnay et al, 2013)."

New references:

Wang, Pinya, et al. "The sensitivity to initial soil moisture for three severe cases of heat waves over Eastern China." *Frontiers in Environmental Science* 7 (2019): 18.  
<https://doi.org/10.3389/fenvs.2019.00018>

De Rosnay, Patricia, et al. "A simplified Extended Kalman Filter for the global operational soil moisture analysis at ECMWF." *Quarterly Journal of the Royal Meteorological Society* 139.674 (2013): 1199-1213. <https://doi.org/10.1002/qj.2023>

pg 13 Table 1, please provide explanation for name acronyms here for easier reference.

We extended Table 1 in order to ease the reference.

| Name | L <sub>De</sub> | g <sub>s</sub> | <u>Description</u>                    |
|------|-----------------|----------------|---------------------------------------|
| STD  | Std             | Std            | <u>Current conditions</u>             |
| LA0  | 0               | Std            | <u>No street trees</u>                |
| LA+  | +50%            | Std            | <u>Increased leaf area density</u>    |
| LA-  | -50%            | Std            | <u>Decreased leaf area density</u>    |
| SC+  | Std             | +50%           | <u>Increased stomatal conductance</u> |
| SC-  | Std             | -50%           | <u>Decreased stomatal conductance</u> |

pg 18 In 4. I didn't find the attribution of the night-time overestimation of air temperature to the inability of COSMO to represent nocturnal stable boundary layer conditions convincing. If it were an

issue with COSMO failing to represent a stable boundary layer, then it should occur in the non-urban site (BLER) night-time cooling rate. However, the supplementary Figure 1 indicates the rate of cooling matches observations well (al-though with a positive bias throughout), which to me indicates the evolution of stability conditions are probably reasonable at the rural site. It therefore seems the overestimation of night-time urban air temperatures could just as easily be from issues within the urban scheme configuration rather than the atmospheric model or stability issues. The supplementary Figure 2 indicates the bulk albedo may be low, and the material properties of facets listed in Table 2 store a lot of heat. I understand those values are based on previous studies, but those values were derived through optimising WRF SLUCM and so are not necessarily realistic urban parameters, but could simply be accounting for deficiencies in that model. Therefore, the parameters won't necessarily be appropriate for BEP-TREE at Basel. I'm not asking for the simulations to be redone, but the authors shouldn't be so quick to simply attribute night-time errors to the atmospheric model when urban parameters may be the issue (low albedo, high heat storage).

We thank the reviewer for pointing out this important aspect.

By re-examining the figures, we agree with the reviewer that the attribution of the night-time overestimation of air temperature at the BKLI tower site (urban site) cannot be explained by the model inability to representing the stable boundary layer.

We agree that the bias may be due to the choice of material properties, especially the too low albedo (as substantiated by the radiation fluxes – Supplement, Figure 2).

However, we remark that some aspects are still inconsistent with the observations, such as the fact that a clear bias was not present at the 2-m stations.

In view of these arguments, we edited the paragraph as follows:

pg 18 ln 1: “The model simulates the evolution of air temperature very well during the evaluation period (Figure 5a) although a slight overestimation with a MBE of 0.51 K and a systematic RMSE of 0.51 K was found. The overestimation occurs mostly during night-time and it seems to be related to an underestimation of the bulk albedo (Supplement) which can be traced back to the choice of material properties (Table 2). A too low albedo may have produced an excess in heat storage with consequently larger sensible heat release at night. Nevertheless, a night-time overestimation of temperature was not found at the near-surface sites, indicating that more analyses are needed to better understand eventual issues with the choice of material properties.”, ~~an issue that was observed also in other urban climate studies with COSMO (Schubert and Grossman-Clarke, 2014; Wouters et al., 2016) and other models (e.g. Lee et al., 2011). The night-time overestimation of air temperature is attributed to the inability of COSMO to represent very stable boundary layer conditions (Cerenzia, 2017).~~”

Additionally, we added a line about this issue in the “Future Work” section.

Pg 27 ln 16: “Even if generally comparable with other models, the performance of COSMO-BEP-Tree still has room for improvements. The model evaluation against urban flux tower measurements revealed a systematic underestimation of specific humidity (q). The source of this bias needs to be



better investigated by, for instance, (a) analysing the soil moisture content and (b) evaluating  $q$  at the model boundaries provided by another model. The model evaluation reveals an overestimation of night-time air temperature above the urban canopy layer, which can be due to the use of default values of material properties. ...

pg 20 ln 27 "are" -> area

Thank you for pointing out the typo.

pg 20 ln 27: "... and leaf area index from street trees ..."

pg 26 "Future Work" It was noted earlier that BEP-TREE does not interact with LSM soil moisture - this is a major limitation which should be discussed here.

We extended the section "Future Work" accordingly.

pg26, ln 18: "In order to improve the representation of the transpiration from street trees and therefore the modelling of the associated latent heat fluxes, a mechanistic stomata model (e.g. Damour et al. (2010)) will have to be implemented. Furthermore, to properly represent soil water scarcity during extended drought periods, an urban hydrology model would have to be implemented (e.g. Järvi 2011, Yang 2015, Stavropoulos-Laffaille 2018)."

#### References:

Järvi, Leena, C. S. B. Grimmond, and Andreas Christen. "The surface urban energy and water balance scheme (SUEWS): Evaluation in Los Angeles and Vancouver." *Journal of hydrology* 411.3-4 (2011): 219-237. <https://doi.org/10.1016/j.jhydrol.2011.10.001>

Yang, Jiachuan, et al. "Enhancing hydrologic modelling in the coupled weather research and forecasting-urban modelling system." *Boundary-Layer Meteorology* 155.1 (2015): 87-109. <https://doi.org/10.1007/s10546-014-9991-6>

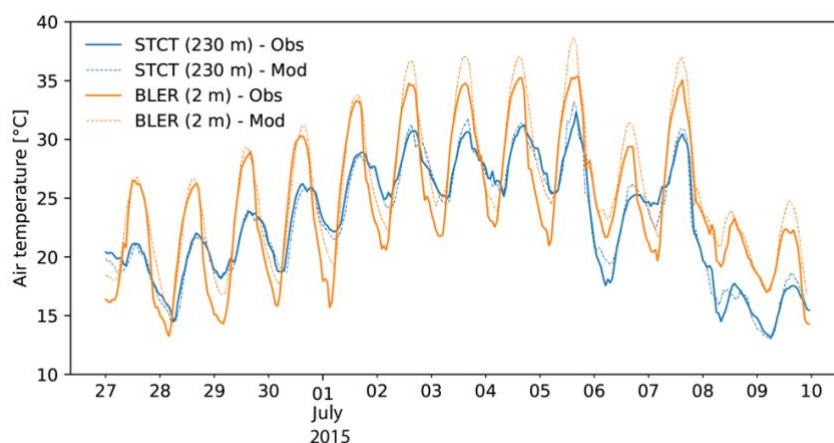
Stavropoulos-Laffaille, Xenia, et al. "Improvements to the hydrological processes of the Town Energy Balance model (TEB-Veg, SURFEX v7. 3) for urban modelling and impact assessment." *Geoscientific Model Development* 11.10 (2018): 4175-4194. <https://doi.org/10.5194/gmd-11-4175-2018>

pg 27 ln 21 "attributed to the inability of the model to reproduce a very stable atmospheric conditions" again, if this argument is used it should be better supported with information from the current study. For example, was the minimum value of turbulent diffusion coefficients ( $K_{i,min}$ ) set at 1 m<sup>2</sup>/s (per Buzzi 2011) in the current experiments? What were the stability conditions at the rural site? Did an inversion form? Why is the rate of cooling at night at the rural site seemingly correct if a stability is incorrectly simulated? Do the observed wind speeds support the conclusion that very stable atmospheric conditions should have formed on those nights? Is there other observational

support for the stability or non-stability of rural sites in the area (non-urbanflux towers, boundary layer measurements etc)?

We thank the reviewer for pointing out this aspect.

We included an additional figure (Figure R2, added to the Supplement) that supports our argument regarding the formation of stable boundary layer conditions. By comparing the air temperatures at the rural reference site BLER (2 m a.g.) and at neighbouring St. Chrischona Tower (STCT, 230 m above ground, 3.4 km away from BLER), we can see that stable boundary layer conditions (characterised by a strong temperature inversion) were present during several nights.

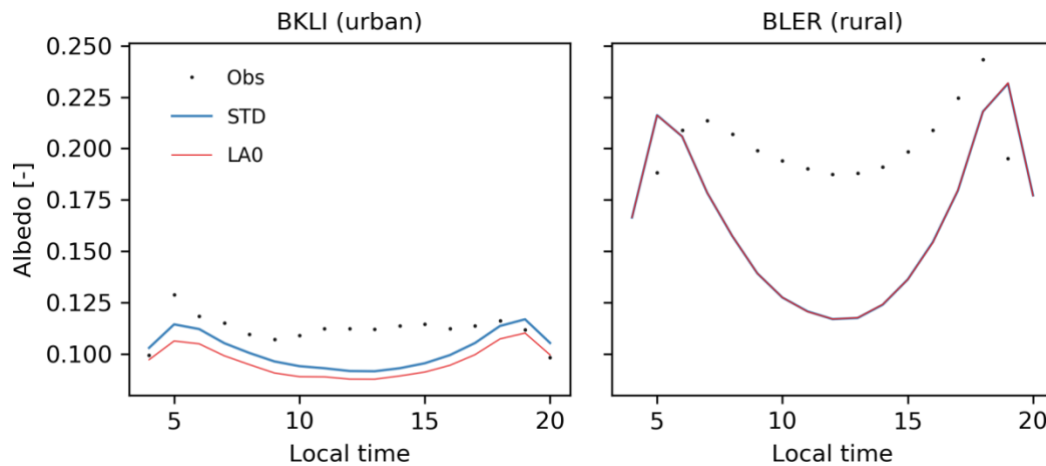


**Figure R2 - Air temperature evolution at the sites of St. Chrischona (STCT, 230 m a.g., lat=47.571767, lon=7.687094) and BLER (2 m a.g.). Solid and dashed lines indicate observations (Obs) and model simulation (Mod), respectively.**

Additionally, we can further see that the night-time air temperature at the rural site of BLER is overestimated, especially during nights with stable boundary layer conditions. By contrast, the model compares very well against the observation at the tower.

However, we agree with the reviewer that a misrepresentation of the stable boundary layer conditions in COSMO cannot fully explain the bias. This is especially clear in the central part of the period, where a positive bias is present throughout the day.

We further examined the data and found a large underestimation of albedo at the rural site (see Figure R3 below, added to the Supplement). For completeness, we included a new figure (in the Supplement) with all the components of the radiation balance at the rural site BLER.



**Figure R3 - Average diurnal profile of observed (Obs) and modelled (STD, LA0) albedo. Albedo is calculated as the ratio between upward and downward short-wave radiation. The site BKLI is characterized by an urban fraction ( $f_{urb}$ ) of 0.79.**

This large underestimation has been already found in another urban study with COSMO-CLM (Schubert et al., 2014) and it is attributed to a misrepresentation of albedo in the land-surface model of COSMO-CLM.

By default in COSMO-CLM, albedo values are determined as a function of soil type, soil moisture and plant fraction. However, there is no distinction between vegetation type and a constant background albedo value for vegetation of 0.15 is applied (Tölle et al., 2018).

For the configurations at the BLER site, the model estimates a daily average value of about 0.15. This value is somewhat smaller than what expected for the land cover at the site, where agricultural fields with crops are present (albedo = 0.18-0.25, Oke et al., 2018). It is not clear, however, whether this bias is a specific problem at the BLER site or a more general issue over the rural areas that surround the city.

Based on this analysis, we recommend that future studies consider the use of more advanced method for albedo estimation, as developed and tested in recent studies with COSMO-CLM (e.g. Tölle et al., 2018).

In conclusion, we think that both aspects (inability to represent very stable boundary layer conditions by COSMO and underestimation of albedo) may play a role in determining the air temperature bias at the rural reference site BLER.

We reconsidered our analysis and made changes to the manuscript as follows:

pg 20 ln 23 “This is probably determined by the combined effect of underestimation in surface albedo and misrepresentation of atmospheric stability (Supplement). A similar underestimation of albedo at rural sites has been already found by Schubert and Grossman-Clarke (2014). The underestimation is attributed to the default albedo scheme of COSMO-CLM, which fails to represent different vegetation types. It is unclear, however, whether this bias is due to specific conditions at the BLER site (grassland) or due to a general misrepresentation in the land-surface model of COSMO-

CLM. The misrepresentation of stable boundary layer conditions is a well-known issue of COSMO (Buzzi et al., 2011; Cerenzia, 2017) and has already been reported in previous studies (Schubert and Grossman-Clarke, 2014; Mussetti et al., 2019)."

pg 27, ln 19-23 "The model evaluation at 2-m reveals an overestimation of the air temperature at the rural sites especially at night. Beside previous studies where this bias was attributed only to the misrepresentation of stable boundary layer conditions (Mussetti et al., 2019), the present study provides new evidences that point out to an underestimation of albedo over rural areas. Future studies need to address this issue that partially limits the model ability to represent the urban heat island effect. Potential solutions include the use of more advanced representations of albedo, already available as options in COSMO-CLM, with explicit consideration of vegetation type or satellite-based albedo values (Tölle et al., 2018). This issue has been already found in other studies (e.g. Schubert and Grossman-Clarke, 2014) and is attributed to the inability of the model to reproduce very stable atmospheric conditions (Buzzi et al., 2011). In respect to the representation of stable boundary layer conditions, more recent versions of the COSMO model (from version 5.4a onward) promise a better performance thanks to a revised turbulence scheme."

#### References:

Schubert, Sebastian, and S. Grossman-Clarke. "Evaluation of the coupled COSMO-CLM/DCEP model with observations from BUBBLE." Quarterly Journal of the Royal Meteorological Society 140.685 (2014): 2465-2483. <https://doi.org/10.1002/qj.2311>

Tölle MH, Breil M, Radtke K and Panitz H-J (2018) Sensitivity of European Temperature to Albedo Parameterization in the Regional Climate Model COSMO-CLM Linked to Extreme Land Use Changes. Front. Environ. Sci. 6:123. <https://doi.org/10.3389/fenvs.2018.00123>

Finally, I found myself referring to the rural figure often throughout the manuscript, it may be better placed in the main body near figure 6.

We added the daily evolution of  $T_{2m}$  at the rural reference site BLER to Figure 6.

pg 21, caption Figure 6: "Figure 6. Comparison between period averaged daily profiles of observed (Obs) and model-simulated 2-m air temperature at the sites BFEL, BSJO, SLFR, ~~and~~ BBIN and BLER during the selected period (discarding the first 5 days as spin-up). Black dots indicate the observations. Blue lines and red lines indicate the results from the STD and LA0 model runs, respectively. The range of variability within the selected period is shown as shaded area for Obs (grey) and STD (light blue)."

pg 28 ln 16 "additionally explained by an underestimation of the night-time temperature" I believe the night-time temperature is overestimated at the BKLI site.

We thank the reviewer for pointing out the typo.

pg 28 ln 16: "At the most urban site, it is additionally explained by an ~~over~~underestimation of the night-time temperature."

# COSMO-BEP-Tree v1.0: a coupled urban climate model with explicit representation of street trees

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**Abstract.** Street trees are more and more regarded as an effective measure to reduce excessive heat in urban areas. However, the vast majority of mesoscale urban climate models do not represent street trees in an explicit manner and for example do not take the important effect of shading by trees into account. In addition, urban canopy models that take interactions of trees and urban fabrics directly into account are usually limited to the street or neighbourhood scale and, hence, cannot be used to analyse the citywide effect of urban greening. In order to represent the interactions between street trees, urban elements and the atmosphere in realistic regional weather and climate simulations, we coupled the vegetated urban canopy model BEP-Tree and the mesoscale weather and climate model COSMO. The performance and applicability of the coupled model, named COSMO-BEP-Tree, are demonstrated over the urban area of Basel, Switzerland, during the heatwave event of June-July 2015. Overall, the model compared well with measurements of individual components of the surface energy balance and with air and surface temperatures obtained from a flux tower, surface stations and satellites. Deficiencies were identified for night-time air temperature and humidity, which can mainly be traced back to limitations in the simulation of the night-time stable boundary layer in COSMO. The representation of street trees in the coupled model generally improved the agreement with observations. Street trees produced large changes in simulated sensible and latent heat flux, and wind speed. Within the canopy layer, the presence of street trees resulted in a slight reduction in daytime air temperature and a very minor increase in night-time air temperature. The model was found to realistically respond to changes in the parameters defining the street trees: leaf area density and stomatal conductance. Overall, COSMO-BEP-Tree demonstrated the potential of (a) enabling city-wide studies on the cooling potential of street trees and (b) further enhancing the modelling capabilities and performance in urban climate modelling studies.

## 1 Introduction

Street trees - trees located in street canyons - and urban vegetation in general are more and more regarded as an effective measure to reduce excessive heat conditions in urban areas (e.g. Shashua-Bar et al., 2009; Armson et al., 2012; Coutts et al.,

2016; Tan et al., 2016; Gunawardena et al., 2017; Manickathan et al., 2018). Excessive heat conditions are typically caused by the combination of regional scale hot weather (heat wave) situations and the urban heat island effect (UHI). The UHI refers to air temperature increase in urban areas compared to their rural surroundings and it is caused by the alteration of the surface energy balance due to the presence of man-made structures and activities (Oke et al., 2017). Climate change is expected to  
5 further accentuate the magnitude and frequency of excessive heat conditions in cities with potentially severe impacts on human health, energy consumption, air pollution and urban ecology (Rosenzweig et al., 2018).

Street trees impact the urban atmosphere through radiative, heat, moisture and flow interactions (Oke, 1989). Street trees intercept radiation through their foliage canopy. By intercepting the direct short-wave (solar) radiation, they reduce the temperature of underlying surfaces (Armson et al., 2012). Additionally, street trees also intercept radiation coming from the canyon  
10 surfaces, such as the long-wave radiation emitted by warm wall and street surfaces. At night, this effect can actually reduce the radiative cooling of wall and street surfaces (Bowler et al., 2010). ~~In terms of heat interactions, tree transpiration reduces the surface temperature of the foliage by converting part of the solar radiation to latent heat~~ Through tree transpiration, leaf surface temperature is reduced resulting in the extraction of sensible heat from the air (Green, 1993). Finally, the tree foliage ~~excerpts~~ extracts momentum from the mean flow, contributing to Turbulent Kinetic Energy (TKE) generation due to wake production  
15 and TKE dissipation due to the small scale of the leaves (Wilson and Shaw, 1977).

~~Modelling studies on the cooling potential of street trees have almost exclusively been performed at~~ The climatic impact of urban vegetation has been investigated in numerous previous studies from the scale of ~~individual street canyons to a single neighbourhood (e.g. Gromke et al., 2015; Ng et al., 2012). Only very few studies investigated the city-wide impact of street trees, mainly due to the limited availability of models able to represent street trees at the scale of the city or at even larger scales.~~  
20 ~~In fact, the vast majority of meso-scale urban climate models only represent vegetation outside the~~ the single street canyon to that of the entire urban region (e.g. Gromke et al., 2015; Ng et al., 2012; De Munck et al., 2018). However, studies on entire urban regions primarily focused on low vegetation, representing low height gardens and green roofs (Wang et al., 2013; De Munck et al., 2013). Street trees, instead, have generally only been represented in a separate natural tile (e.g. Schubert and Grossman-Clarke, 2013; Li and Norford, 2013). This approach precluded considering any interactions between trees and urban surfaces in a street canyon, neglecting important  
25 ~~effects such as the shading effect of trees on the canyon's surfaces such as shading and sheltering effects (Kravenhoff et al., 2014, 2015). Other studies employed somewhat more sophisticated methods but still neglect some of the critical interactions between trees, canyon surface and airflow (Thatcher and Hurley, 2012).~~

The very first example of a weather and climate model with an integrated representation of street trees was developed by Dupont et al. (2004). They built an urban canopy model, called DA-SM2-U, for the Penn State-NCAR fifth-generation  
30 Mesoscale Model (MM5, Grell et al. (1994)). DA-SM2-U was based on the multi-layer urban canopy concept where a drag-force approach is used to represent the dynamic and turbulent effects of buildings and vegetation. In the meantime, the Weather and Forecasting Model (WRF, Skamarock et al. (2005)) has replaced MM5 and, to the authors best knowledge, DA-SM2-U was not ported to WRF. Lee and Park (2008) developed an urban canopy model with an explicit representation of street trees named Vegetated Urban Canopy Model (VUCM). VUCM adopts a single-layer urban canopy concept (i.e., the urban canopy  
35 is represented by only one model layer) where street trees are represented as a single tree canopy characterised by canopy

cover fraction, mean height and mean leaf area density. Lee (2011) further developed VUCM by including a grass-covered soil surface within the canyon. Lee et al. (2016) incorporated VUCM into WRF and performed an application over Seoul, South Korea. VUCM was also used by Loughner et al. (2012), coupled with WRF, to simulate the role of street trees in Washington, D.C., USA. Wang et al. (2018) incorporated a representation of the radiative effects of street trees (Wang, 2014) into the single-layer urban canopy model of WRF (Chen et al., 2011a). They simulated the impact of street trees over the contiguous United States using a 20 km model resolution. Krayenhoff et al. (2018) used WRF, coupled with the single-layer urban canopy model, to simulate the impact of urban adaptation strategies (including street ~~trees~~-tree planting) for air temperature during contemporary and future climate over the contiguous United States. A couple of other urban canopy models (UCMs) with explicit representation of street trees have been developed (Ryu et al., 2016; Redon et al., 2017) but, to the authors best knowledge, not yet coupled with a mesoscale weather and climate model.

The literature review revealed a number of major limitations in the models used to study the city-wide impact of street trees. Several studies neglected some the interactions between street trees and urban fabrics. For example, the exchange of long-wave radiation between street trees and urban fabrics is often neglected (Wang et al., 2018; Krayenhoff et al., 2018), together with the impact of street trees on the flow field. Neglecting long-wave radiation exchanges and impacts of trees on the flow field may overestimate their cooling potential. Another simplistic assumption ~~regarded~~regards ~~concerns~~concerns the representation of tree shading, which was in some cases independent of the solar direction (Dupont et al., 2004; Lee and Park, 2008).

A second common assumption ~~regards~~regards ~~concerns~~concerns the characterisation of street trees in model applications. In all the previous studies, no real data on the location, density height and species of street trees were used. Instead, a homogeneous distribution of street trees has often been assumed (Dupont et al., 2004; Loughner et al., 2012). The increasing availability of detailed street ~~trees~~-tree datasets (e.g. Alonzo et al., 2014; Konarska et al., 2016) gives now a chance to represent the distribution of street ~~trees~~-tree characteristics more realistically during model applications.

Finally, the majority of the studies did not include a comprehensive evaluation of their model. Typically, only the underlying urban canopy model used in the coupled model has been evaluated off-line. Online model evaluation against surface stations, flux towers and satellite observations ~~-,in-the-presence-of-street-trees,-~~ are needed to better understand the model performance and to identify future lines of development.

Motivated by the great potential of using an integrated model to study the urban climate impact of street trees in a city-wide manner, the objective of this study is two-fold. First, we aim to document the development of an urban climate model with explicit representation of street trees, featuring a two-way coupling between the vegetated urban canopy model BEP-Tree (Krayenhoff, 2014; Krayenhoff et al., 2019) and the mesoscale weather and climate model COSMO-CLM (COSMO hereinafter; Rockel et al. (2008)). The coupled model, named COSMO-BEP-Tree, represents the multiple interactions between street trees, urban fabrics and the atmosphere in a comprehensive way considering the exchange of heat, moisture, and momentum and the transfer of short- and long-wave radiation. COSMO-BEP-Tree enables city-wide studies on the cooling potential of street trees in comparison, or in combination, with other UHI mitigation measures (e.g. albedo changes). Additionally, it has the potential to enhance the ~~modelling~~ performance in urban climate modelling studies by a more realistic representation of the



surface and its interaction with the atmosphere. In order to facilitate the application, the model comes with a tool to pre-process the additional inputs required for COSMO-BEP-Tree, such as street ~~trees-tree~~ datasets or a 3D city model.

Second, we present an application and comprehensive evaluation of COSMO-BEP-Tree over the transnational urban agglomeration of Basel (Switzerland, Germany and France) during a heatwave event in June-July 2015. The evaluation makes use of the extensive measurement infrastructure for urban climate studies available in Basel (Feigenwinter et al., 2018; Wicki et al., 2018) including an urban flux tower and a network of surface stations and additionally uses land surface temperature (LST) observations from satellite. In order to verify the model's response to the parameters that define the street trees, a sensitivity analysis is also presented.

We remark that this study only aims to document the development and evaluation of COSMO-BEP-Tree. Analysis on the role of street trees in urban climate and on their potential for mitigating excessive heat conditions are out of the scope of this paper and will be presented in follow-up studies. Although the focus of this study is on street trees, it is important to consider that other components of the so-called urban forest, such as trees in parks and private ~~garden-gardens~~ as well as peri-urban trees, are also likely to have a remarkable impact on the urban climate.

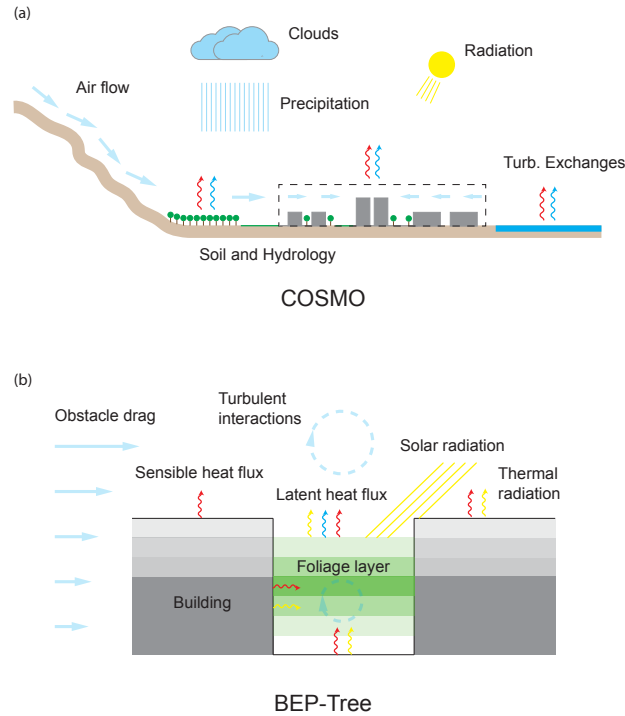
The manuscript is structured as follows. The two model components COSMO and BEP-Tree, together with the coupling strategy and the pre-processor, are described in Section 2. Section 3 ~~presents-and~~ 4 present the model evaluation and sensitivity. In Sections 5 and 6 future work and conclusions are discussed.

## 2 Model description

### 2.1 COSMO model

The Consortium for Small-scale Modelling (COSMO) model is a non-hydrostatic limited-area atmospheric prediction model (COSMO, 2019). It has been designed for operational numerical weather prediction, regional climate simulation and other applications at the mesoscale (weather phenomena of sizes between about 1 km and a few hundred kilometres). COSMO evolved from the operational weather forecast Local Model (LM) of the German Weather Service (Steppeler et al., 2003) and has been developed by a consortium of weather services in Europe and by the CLM-Community for climate applications (CCLM-Community). The regional climate model version of COSMO, called COSMO-CLM, includes modifications allowing the application on time-scales up to centuries (Rockel et al., 2008). These modifications comprise a representation of phenological cycles as well as externally prescribed, time-dependent atmospheric CO<sub>2</sub> concentrations.

The COSMO model is based on the thermo-hydrodynamical equations describing non-hydrostatic compressible flow in a moist atmosphere. A variety of physical processes are taken into account by parameterisation schemes: ~~sub-grid-scale turbulence, surface fluxes, grid and sub-grid-scale clouds and~~ cloud formation and precipitation, moist and shallow convection, ~~radiation, soil and vegetation, lake and sea ice schemes~~ sub-grid scale turbulence, radiation and land-surface exchanges (see Figure 1a). At the convection-resolving model resolution used in this study, the parameterisations for sub-grid scale clouds and moist convection are switched off. Additional information on the governing equations and physical parameterisations can be found in the model documentation (COSMO, 2019).



**Figure 1.** Schematic of the key processes represented by (a) COSMO and (b) BEP-Tree for urban climate simulations. Details about the interactions between radiation and the foliage layers (e.g. radiation absorption and scattering) are not represented here. Please refer to the Sections 2.1 and 2.2 for a comprehensive description of the models.

Recently, urban parameterisations of different degrees of complexity have been developed for and coupled with the COSMO model (Trusilova et al., 2016). The urban parameterisation TERRA-URB (Wouters et al., 2015, 2016) employs a simple but efficient approach by representing the urban canopy using a limited number of parameters (so-called bulk approach). The Town Energy Balance (TEB) implementation of Trusilova et al. (2013) represents an intermediate level of complexity with a single-layer urban canopy model. Finally, the multi-layer urban canopy model DCEP (Double-Canyon Effect parameterisation) employs the most physically-based but input-demanding representation of the urban canopy by solving the radiation exchange in a realistic multi-layer canyon configuration (Schubert et al., 2012).

Applications of these models included studies of the impact of climate change on cities (Grossman-Clarke et al., 2017; Wouters et al., 2017), the influence of green areas and low albedo materials (Schubert and Grossman-Clarke, 2013) and the intra-urban climate variability (Mussetti et al., 2020).

## 2.2 Building Effect parameterisation with Trees (BEP-Tree)

The Building Effect Parameterisation with Trees (BEP-Tree) is a multi-layer urban canopy model with explicit representation of trees and their interaction with the street canyon (Krayenhoff, 2014; Krayenhoff et al., 2019). Building on the geometry

of the existing multi-layer urban canopy model BEP (Martilli et al., 2002), Krayenhoff et al. (2014) included the effects of tree foliage on the radiative exchange within the street canyon.

The basic geometry is a two-dimensional canyon with vertical leaf area density profiles and probabilistic variation of building height (see Figure 1b). Tree foliage is permitted both between and above buildings allowing the representation of street trees as well as trees overhanging buildings.

Ray tracing is used to determine the receipt of direct short-wave irradiance by building and foliage elements. View factors for long-wave and short-wave diffuse radiation exchange are computed once at the start of the simulation using a Monte Carlo ray tracing approach, and used thereafter to calculate multiple reflections between all urban fabrics. Interception of radiation by tree foliage is modelled with the Bouguer-Lambert-Beer law:

$$\Delta V_i = r_i \left( 1 - \exp(-K \Omega L_D \Delta s f_i) \right), \quad (1)$$

where  $\Delta V_i$  is the reduction in intensity of ray  $i$  due to the tree foliage [ $\text{W m}^{-2}$ ],  $r_i$  is the initial intensity of the ray  $i$  [ $\text{W m}^{-2}$ ],  $K$  is the foliage extinction coefficient [-],  $L_D$  is the leaf area density [ $\text{m}^2 \text{m}^{-3}$ ],  $\Omega$  is the foliage clumping index [-],  $\Delta s$  is the 2-D ray step size [m] and  $f_i$  is the ratio of 3-D (actual) to 2-D (model) distance travelled by the ray (actual 3-D paths of all rays are mapped to the 2-D domain, depending on the canyon orientation; Krayenhoff et al. (2014)).

BEP-Tree includes a parameterisation of building and tree foliage effects on the airflow (Martilli et al., 2002; Santiago and Martilli, 2010; Krayenhoff et al., 2015). The parameterisation introduces (a) a sink term for momentum to account for obstacle drag, (b) modifications of turbulence length scales by building interactions, and (c) enhanced dissipation of turbulent kinetic energy due to the small scale of the tree foliage elements. The obstacle drag from tree foliage is determined as:

$$S_{m,i} = -C_{dv} \Omega L_D U \overline{u_i}, \quad (2)$$

where  $S_{m,i}$  is the sink term for momentum along the coordinate axis  $i$  [ $\text{m s}^{-2}$ ],  $C_{dv}$  is the sectional drag coefficient for tree foliage (set to 0.2) [-],  $U$  is the horizontal wind speed [ $\text{m s}^{-1}$ ] and  $\overline{u_i}$  is the wind speed component along axis  $i$  [ $\text{m s}^{-1}$ ].

Surface energy balance equations for roof, wall and road elements are identical to Martilli et al. (2002). For leaf layers, the Penman equation is rearranged to solve for the leaf temperature as in Campbell and Norman (2012). The sensible heat flux, latent heat flux and source of moisture from street trees are determined as:

$$Q_H = c_p g_{Ha} \Omega L_D (T_{\text{veg}} - T_{\text{air}}), \quad (3)$$

$$Q_E = \lambda \frac{g_s g_a}{g_s + g_a} \Omega L_D \left( s (T_{\text{veg}} - T_{\text{air}}) + \frac{e_s - e_{\text{air}}}{p} \right), \quad (4)$$

$$S_q = \frac{\mu_w}{\rho_{\text{air}}} \frac{g_s g_a}{g_s + g_a} \Omega L_D \left( s (T_{\text{veg}} - T_{\text{air}}) + \frac{e_s - e_{\text{air}}}{p} \right), \quad (5)$$

where  $Q_H$  is the sensible heat flux from foliage per unit height [ $\text{W m}^{-3}$ ],  $c_p$  is the specific heat capacity of air [ $\text{J mol}^{-1} \text{K}^{-1}$ ],  $g_{Ha}$  is the heat conductance [ $\text{mol m}^{-2} \text{s}^{-1}$ ],  $T_{\text{veg}}$  is the temperature of the foliage layer [K],  $T_{\text{air}}$  is the air temperature at the height of the foliage layer [K],  $Q_E$  is the latent heat flux from foliage per unit height [ $\text{W m}^{-3}$ ],  $\lambda$  is the latent heat of vaporisation [ $\text{J mol}^{-1}$ ],  $g_s$  is the vapour stomatal conductance of the leaf [ $\text{mol m}^{-2} \text{s}^{-1}$ ],  $g_a$  is the vapour boundary-layer

conductance of the leaf [ $\text{mol m}^{-2} \text{s}^{-1}$ ],  ~~$\rho_{\text{air}}$  is the density of air  $\text{kg m}^{-3}$~~ ,  $s = \frac{de_s}{dT_{\text{air}}} \frac{1}{p}$  is the change in saturation pressure with temperature given by the Clausius-Clapeyron equation (Stull, 2012) [ $\text{K}^{-1}$ ],  $e_s$  is the saturation vapour pressure [Pa],  $e_{\text{air}}$  is the vapour pressure in the air [Pa],  $p$  is the atmospheric pressure [Pa],  $S_q$  is the source term for moisture due to transpiration from foliage [ $\text{s}^{-1}$ ]~~and~~,  $\mu_w$  is the molar mass of water [ $\text{kg mol}^{-1}$ ] ~~and  $\rho_{\text{air}}$  is the density of air [ $\text{kg m}^{-3}$ ]~~. Here we report

5 only those equations that are relevant for the coupling of BEP-Tree with COSMO and for understanding the response to the sensitivity experiments described later. The complete set of equations as well as further details about the model components can be found in Krayenhoff et al. (2014, 2015, 2019), including a recently-developed parameterisation for climate impacts of street tree foliage distribution.

The main model inputs to BEP-Tree are (a) atmospheric state variables above the canyon and (b) canyon geometry including

10 vegetation. The former includes the traditional atmospheric variables and radiation fluxes, that can be either provided by a tall tower (off-line application) or by a weather model (online application). The ~~latter includes~~ canyon geometry inputs include height distribution of buildings, canyon width, thermo-physical material properties of active surfaces (wall elements, roof and street) and parameters representing in-canyon vegetation.

Regarding in-canyon vegetation, a vertical profile of leaf area density ( $L_D$ ) needs to be provided for both the canyon and the

15 building columns.  $L_{D,\text{can}}$  and  $L_{D,\text{bld}}$  represent the surface area of leaves per unit volume of air [ $\text{m}^2 \text{m}^{-3}$ ] in the canyon space and in the building space (e.g. in case of trees overhanging buildings), respectively. Additionally, information on the spatial distribution of the leaves from street trees in the available canyon space needs to be provided. This information is represented by the clumping index  $\Omega$  as defined by Nilson (1971). The index takes values between 0 and 1, with  $\Omega = 1$  being the case where the leaves are randomly distributed.  $\Omega$  encompasses clumping at several scales, from the within-branch to the between-crown

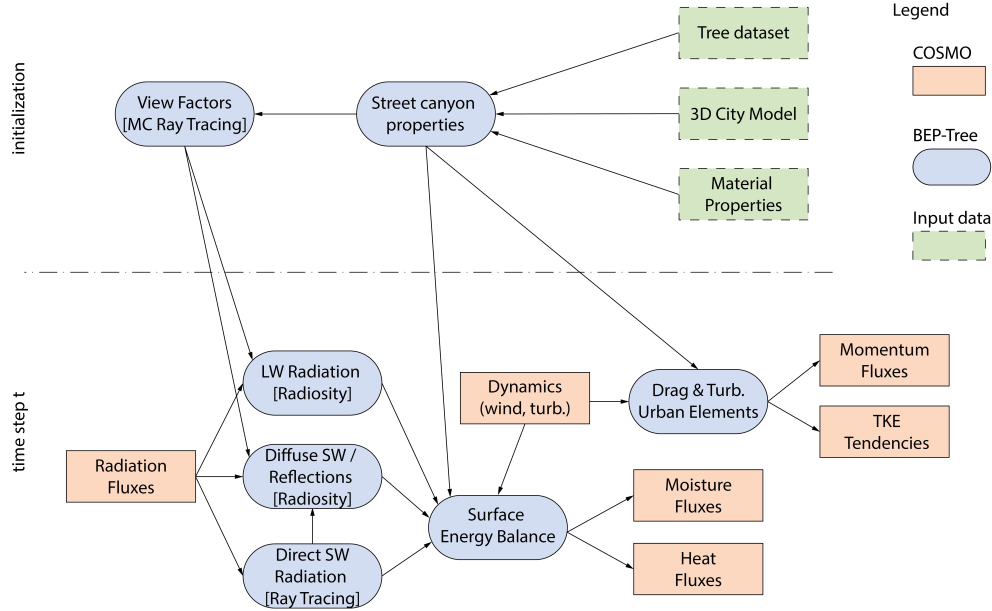
20 scale.  $\Omega$  is used, together with  $L_D$ , to account for the fraction of leaves that actively contribute to the radiation exchange. Moreover,  $\Omega$  is used to represent the aerodynamic sheltering following the approach proposed by Marcolla et al. (2003) and to modulate wake production and turbulent kinetic energy (TKE) dissipation (Krayenhoff, 2014).  $\Omega$  is always used in combination with  $L_D$  to compute an “effective leaf area density” defined as  $L_{De} = L_D \Omega$ .

BEP-Tree has been evaluated in off-line mode against ~~a measurement site~~ measurements in Vancouver, Canada (Krayenhoff,

25 2014). The evaluation made use of measured fluxes of radiation and turbulent exchange in the inertial sublayer, as well as measurements of air temperature and humidity in the urban canopy layer. The evaluation has been recently extended using measurements of street surface temperature and street-level radiation fluxes from two sites in North America (Krayenhoff et al., 2019).

### 2.3 COSMO-BEP-Tree

30 We coupled BEP-Tree to the COSMO-CLM model v5.0\_clm2.1. The coupling is done through model integration (also called joint coupling by Brandmeyer and Karimi (2000)) and makes use of the interface created by Schubert et al. (2012) for the urban parameterisation DCEP. ~~The coupled model, named COSMO-BEP-Tree, is developed to simulate the exchange of energy, momentum, and moisture between an urban canopy with street trees and the atmosphere and the effects of buildings and trees on TKE during weather and climate simulations.~~ The coupling makes use of the tile approach, where mixed ~~configurations~~



**Figure 2.** Flow chart of the coupling of BEP-Tree with COSMO. Rose boxes denote operations-quantities of COSMO, blue boxes of BEP-Tree. Additional input datasets required to run COSMO-BEP-Tree are shown as green boxes.

proportions of urban and natural surfaces within a grid cell can be considered. Natural surfaces within a urban grid cell represents, for instance, urban parks and peri-urban forests. The surface-atmosphere interactions are calculated independently for the urban and rural tiles and then averaged with respect to the fractional coverage of each tile. BEP-Tree is used to calculate the surface-atmosphere interactions for the urban tile, while the land surface model of COSMO does it for the natural tile (see Section 2.1 for further details).

A schematic representation of the coupling is presented in Fig. 2. Before executing COSMO-BEP-Tree, the parameters that characterise the street canyon are estimated for all urban grid cells in the model domain from external input datasets of trees, building geometries and imperviousness. This task is performed by a dedicated pre-processor described in Sec. 2.4. Once all required inputs are generated, the coupled model can be executed. At the first model time step, the urban canopy parameters are used to estimate the view factors for long-wave and diffuse short-wave radiation between each street canyon element (sky; wall, roof, street and foliage layer) including the sky. The view factors are stored in the form of a matrix and applied during the model simulation at each time step. In the case of direct short-wave radiation, the receipt of radiation by each street canyon element is computed online during the simulation using a direct ray tracing algorithm. Subsequently, the diffuse short-wave radiation, short-wave reflections and received and emitted long-wave radiation at each street canyon element are estimated using a radiosity approach (making use of the view factor matrices). At this point, the radiation balance at each active surface is available and is used to determine the radiation balance of the entire grid cell in the COSMO model.

The user can specify the frequency at which the radiation exchanges in BEP-Tree are updated. By default, this is set to 0.05 h (3 minutes), as a compromise between accuracy and computational cost.

The air temperature tendencies (increments) due to sensible heat flux, the momentum tendencies due to obstacle drag and the TKE tendencies are determined for each street canyon element. For the foliage layers, latent heat and moisture fluxes are calculated by solving the leaf energy balance. Although moisture exchange between street trees and the atmosphere is implemented, street trees do not interact with soil moisture content as represented by COSMO's land surface scheme. In other words, a mechanistic interaction between soil moisture and the transpiration of street trees is not included, assuming that the transpiration is never limited by soil water availability. A careful representation of soil moisture in the urban tiles would have required a new urban hydrology scheme, which was beyond the scope of the study. The missing interaction with soil moisture may reduce the model ability to represent variations in transpiration during periods with large changes in soil moisture (Konarska et al., 2016; Asawa et al., 2017). Nevertheless, street trees are less sensitive to variations in soil moisture than short vegetation, thanks to their deeper root system (Chen et al., 2011b; Asawa et al., 2017).

The radiative coupling between COSMO and BEP-Tree is realised through four bulk radiation parameters: (a) albedo for diffuse short-wave radiation  $\alpha_{\text{urb}}^{\downarrow}$ , (b) albedo for direct short-wave radiation  $\alpha_{\text{urb}}^{\downarrow\downarrow}$ , (c) mean emissivity  $\epsilon_{\text{urb}}$  and (d) radiative surface temperature  $T_{\text{urb}}$ .  $\alpha_{\text{urb}}^{\downarrow}$  does not change with time being only dependent on the canyon geometry and on the albedo of individual surface elements, given the assumption of isotropic diffuse radiation exchange. Therefore, it is calculated only once at the beginning of the simulation as:

$$\alpha_{\text{urb}}^{\downarrow} = \frac{K_{\text{urb}}^{\uparrow 0}}{K^{\downarrow}}, \quad (6)$$

with  $K_{\text{urb}}^{\uparrow 0}$  being the shortwave radiation reflected from all street canyon elements into the sky for the case where the incoming direct short-wave radiation  $K^{\downarrow\downarrow}$  is zero and  $K^{\downarrow}$  being the incoming diffuse short-wave radiation.

$\alpha_{\text{urb}}^{\downarrow\downarrow}$ , instead, depends on the position of the sun and is therefore estimated at each time step as:

$$\alpha_{\text{urb}}^{\downarrow\downarrow} = \frac{(K_{\text{urb}}^{\uparrow} - \alpha_{\text{urb}}^{\downarrow} K^{\downarrow})}{K^{\downarrow\downarrow}}, \quad (7)$$

with  $K_{\text{urb}}^{\uparrow}$  being the shortwave radiation reflected from all street canyon elements into the sky.

The radiative surface temperature is calculated as:

$$T_{\text{rad}} = \left( \frac{L_{\text{urb}}^{\uparrow} - (1 - \epsilon_{\text{urb}}) L^{\downarrow}}{\sigma \epsilon_{\text{urb}}} \right)^{1/4}, \quad (8)$$

where  $L_{\text{urb}}^{\uparrow}$  is the sum of the emitted and reflected long-wave radiation from all street canyon elements into the sky,  $\epsilon_{\text{urb}}$  is the bulk emissivity of the urban surface,  $L^{\downarrow}$  is the incoming long-wave radiation and  $\sigma$  is the Stephan-Boltzmann constant.  $\epsilon_{\text{urb}}$  is calculated as the mean emissivity of the street canyon elements.

The bulk radiation parameters from the urban tile ( $_{\text{urb}}$ ) computed by BEP-Tree are combined with the ones computed by COSMO for the natural ( $_{\text{nat}}$ ) tile as:

$$\alpha^{\downarrow} = f_{\text{urb}} \alpha^{\downarrow}_{\text{urb}} + f_{\text{nat}} \alpha^{\downarrow}_{\text{nat}}, \quad (9)$$

$$\alpha^{\uparrow} = f_{\text{urb}} \alpha^{\uparrow}_{\text{urb}} + f_{\text{nat}} \alpha^{\uparrow}_{\text{nat}}, \quad (10)$$

$$5 \quad \epsilon = f_{\text{urb}} \epsilon_{\text{urb}} + f_{\text{nat}} \epsilon_{\text{nat}}, \quad (11)$$

$$T_{\text{rad}} = \left( \frac{(f_{\text{urb}} \epsilon_{\text{urb}} T_{\text{urb}}^4 + f_{\text{nat}} \epsilon_{\text{nat}} T_{\text{nat}}^4)}{\epsilon} \right)^{1/4}, \quad (12)$$

with  $f_{\text{urb}}$  and  $f_{\text{nat}}$  ( $= 1 - f_{\text{urb}}$ ) being the planar fractions of urban and natural surfaces, respectively. The resulting bulk radiation parameters are used as lower (surface) boundary conditions for the short-wave and long-wave radiation routines in COSMO.

In the same way, the tendencies of momentum, sensible heat, moisture and TKE production from the urban tile are combined with the corresponding tendencies from the natural tile.

BEP-Tree operates on a grid (called urban grid) that is vertically independent of and generally finer than the grid of COSMO (called mesoscale grid). As a default, the vertical resolution of the urban grid is 5 m. At each model time step, the input fields from COSMO have to be interpolated to the urban grid whereas the output fields (tendencies) of BEP-Tree have to be aggregated back to the mesoscale grid. The interpolation and aggregation of fields in COSMO-BEP-Tree follows the implementation of Schubert (2013).

~~In analogy to what done by~~ Following Schubert (2013), the 2 m air temperature is diagnosed from the air temperature at the lowest mesoscale model layer and the surface temperature from both urban and natural surfaces as:

$$T_{2\text{m}} = (f_{\text{nat}} T_{\text{g}} + f_{\text{urb}} T_{\text{sstr}}) + r (T_1 - (f_{\text{nat}} T_{\text{sg}} + f_{\text{urb}} T_{\text{gstr}})), \quad (13)$$

where  $T_{\text{g}}$  is the surface temperature of the natural tile,  ~~$T_{\text{s}}$  is the surface temperature of the urban tile~~  $T_{\text{str}}$  is the street surface temperature,  $r$  is a coefficient that represents the stability-dependent resistance for scalar transport as estimated by the TKE-based surface-layer transfer scheme of COSMO (Baldauf et al., 2011), and  $T_1$  is the air temperature at the lowest model level.

## 2.4 COSMO-BEP-Tree preprocessor

Additional input variables need to be calculated before running COSMO-BEP-Tree. These variables, referred to as urban canopy parameters (UCPs) in the following, describe the spatial representation of the street canyon (including street trees) required by the urban canopy model. A dedicated python program, ~~UPCgenerator~~ UCPgenerator v1.0, has been developed to generate the UCPs based on external datasets. The workflow of UCPgenerator is shown in Fig. 2 (green boxes). A technical description of UCPgenerator v1.0 is given in Appendix A.

The UCPgenerator requires the following three input datasets: (1) street trees, (2) buildings and (3) imperviousness. The dataset of street trees consists of a spatial map of the street tree distribution. This may include information on tree height, leaf area index ( $L$ ) or leaf area density ( $L_D$ ) and tree species. In the case of missing information, representative profiles or knowledge-based assumptions can be used instead. Information on street trees is typically provided by inventories maintained

by urban environmental agencies (e.g. Stadtgaertnerei Kantons Basel-Stadt), remotely sensed data such as LIDAR data (e.g., Alonzo et al., 2014) and combined data derived using data fusion techniques (e.g., Branson et al., 2018).

The building dataset consists of a collection of building geometries in the urban area, including building height as an attribute of each individual building. 3D building datasets are typically provided by national mapping agencies and collaborative projects (e.g. OpenStreetMap).

The imperviousness dataset is a spatial representation of the horizontal fraction of impervious surfaces in the model domain. Such data is available at the European scale (e.g. EEA, 2015) as well as at the global scale (e.g. Brown de Colstoun et al., 2017).

### 3 Model evaluation

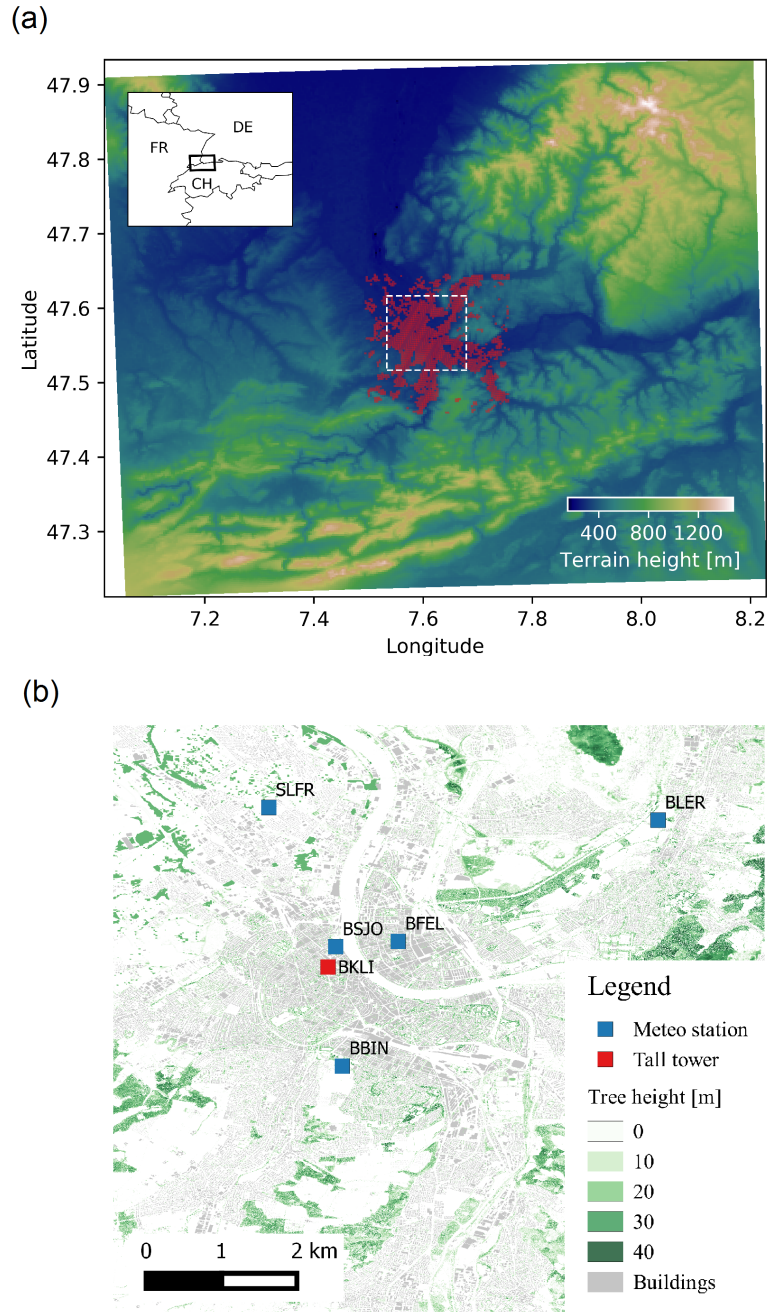
#### 3.1 Study area and model set up

COSMO-BEP-Tree was applied over the Basel metropolitan area (Basel hereinafter), located mainly in Switzerland but extending over the borders of France and Germany (Fig. 3). Situated along the Rhine river, the larger Basel area has a population of approx. 830<sup>+</sup>,000 inhabitants and is surrounded by hilly to mountainous terrain especially towards the south (Jura mountains) and north-east (Black Forest). According to the Köppen climate classification, Basel features a temperate oceanic climate (Köppen: Cfb). The inner city (Basel-Stadt) ~~is characterised by an abundance of green areas with~~ includes more than 24<sup>+</sup>,000 urban trees and 275 hectares of public vegetated surfaces distributed over a total area of the city centre of approx. 2<sup>+</sup>,385 hectars (Stadtgaertnerei Kantons Basel-Stadt). The average canopy cover from street trees is about 20% when considering the entire urban area, but it varies from 6 to 31% in the different neighbourhoods (Supplement). The city of Basel is ~~not only~~ an interesting target for ~~its abundant green infrastructures, but also for~~ the availability of extensive observational data sets for model evaluation, which had been used in numerous previous urban climate studies (e.g. Rotach et al., 2005; Parlow et al., 2014). Observations and input data used here are described in Section 3.3 and Section 3.2.

In this study, COSMO-BEP-Tree was run for the period 22 June 2015 - 9 July 2015, with the first 5 days discarded as spin-up period. The ~~period analyses were performed on the period 27 June 2015 - 9 July 2015, called study period hereinafter.~~ The period corresponds to an intense heatwave that persisted across Europe producing a seasonal mean surface air temperature 2.4 K above the 1964-1993 mean (Dong et al., 2016; Ionita et al., 2017). The event featured a stable anticyclonic weather regime with a prevalence of cloud-free conditions, low wind speeds and absence of precipitation. The conditions were ideal for evaluating the model performance, particularly with respect to the impact of street trees on the urban atmosphere and the development of the UHI.

We applied COSMO-BEP-Tree over a domain of approx. 90 km x 80 km with a horizontal grid spacing of 0.0025° (approx. 270 m). 60 levels were used in the vertical direction, with 5 and 23 levels in the first 100 m and 1000 m, respectively. The model set-up closely follows the configuration described in ~~?~~ Mussetti et al. (2020). Initial and boundary conditions were taken from the operational COSMO-2 analyses operated by the Swiss Federal Office of Meteorology and Climatology (MeteoSwiss), which are centred over Switzerland and cover the whole Alpine domain. The analyses have a horizontal resolution of about





**Figure 3.** (a) The model domain centred over the metropolitan area of Basel. The terrain height (in m a.s.l.) is shown in the background. The urban area is highlighted in red. (b) Detail of the urban area ([corresponding to the dashed rectangle in \(a\)](#)) with the location of the measurement sites used for the evaluation (Section 3). In the background are the tree height distribution (green) (Basel-Stadt, 2019) and building geometries (grey).

**Table 1.** Summary of numerical experiments and changes in the parameters used for the sensitivity study. Std indicates the standard value(s) as provided in Section 3.2. The other parameters are kept constant with the values provided in Section 3.2.

| Name | $L_{De}$ | $g_s$ | Description                           |
|------|----------|-------|---------------------------------------|
| STD  | Std      | Std   | <u>Current conditions</u>             |
| LA0  | 0        | Std   | <u>No street trees</u>                |
| LA+  | +50%     | Std   | <u>Increased leaf area density</u>    |
| LA-  | -50%     | Std   | <u>Decreased leaf area density</u>    |
| SC+  | Std      | +50%  | <u>Increased stomatal conductance</u> |
| SC-  | Std      | -50%  | <u>Decreased stomatal conductance</u> |

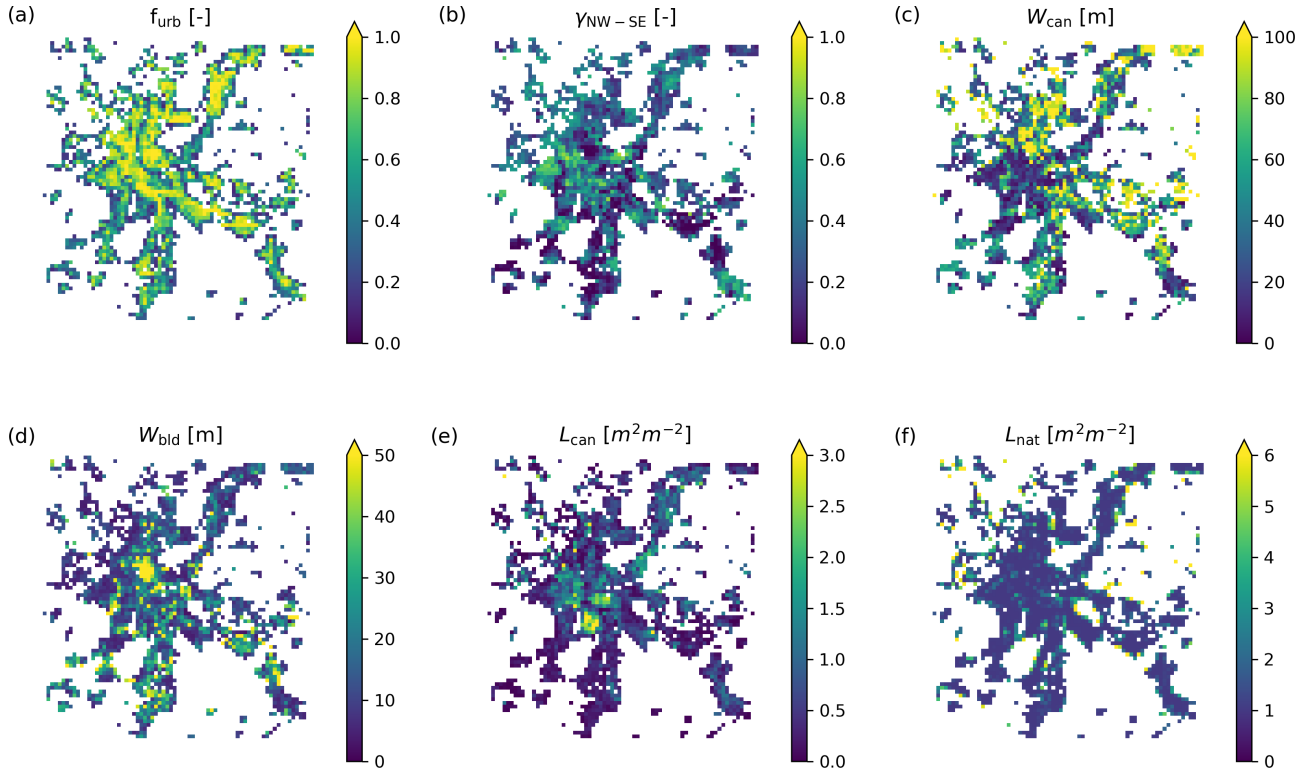
2 km and are produced using a nudging technique (Schraff, 1997) applied to near-surface observations and vertical soundings of pressure, humidity and wind. Soil moisture, a key variable for accurate representation of heatwaves (Wang et al., 2019), is also initialised with the COSMO-2 analyses, which trace back to the soil moisture analyses of ECMWF’s operational IFS model (De Rosnay et al., 2013). As additional input parameters for COSMO, the 1” global digital elevation map ASTER (Tachikawa et al., 2011) for topography, the 10” land use dataset GlobCover 2009 (Loveland et al., 2000) for land cover, and the 30” Harmonized World Soil Database (FAO et al., 2009) for soil properties were used.

To assess the physical consistency of the implementation of the different interactions with street trees, we performed additional sensitivity simulations changing the values of the parameters describing the street trees (Table 1). In addition to the standard run (STD), simulations without street trees (LA0), with increased and decreased values of leaf area density (LA+ and LA-, respectively), and with increased and decreased values of stomatal conductance (SC+ and SC-, respectively) were performed.

### 3.2 Urban canopy input data

Figure 4 shows the distribution of the main urban canopy parameters as derived by the preprocessor UCPgenerator. The model represents the spatial variability of the urban texture in terms of fraction of urban surfaces, street canyon orientation and width , building width and urban vegetation (Figures 4e-f).  $L_{can}$  and  $L_{nat}$  are obtained by vertical integration of the respective leaf area densities  $L_{D,can}$  and  $L_{D,nat}$ . Figure 4 also gives a flavor of the heterogeneity of the city: Industrial areas in the northern and south-east parts of the city, for instance, are associated with wide street canyons and very low vegetation density. Residential areas in the ~~centre-south and centre-west~~south and west of the centre, in contrast, are characterised by narrow street canyons and relatively high vegetation density.

The urban canopy parameters were derived with UCPgenerator from the following input datasets: (a) imperviousness from EEA (2015), (b) 3D buildings from Federal Office of Topography (2007) and (c) lidar-based tree canopy height (see data availability section). The imperviousness dataset represents the percentage of soil sealing (including building area) and has a resolution of 20 m. The building dataset is a vector dataset representing the building geometries (accuracy 3-8 m), including



**Figure 4.** Urban canopy parameters derived for the model domain: (a) urban fraction, (b) canyon fraction along the NW-SE direction, (c) average canyon width, (d) average building width, (e) leaf area index due to street trees, and (f) leaf area index due to vegetation outside the urban canopy. Further details about the variables [are given](#) in Table A1. Only urban grid cells ( $f_{urb} > 0.1$ ) are shown.

[heights but without information on the building materials or use. The](#) tree canopy height dataset has a 2 m spatial resolution and includes all trees in the Basel area. The imperviousness data set was used to differentiate between street trees (contributing to  $L_{D,can}$ ) and trees over natural surfaces (contributing to  $L_{D,nat}$ ). Specifically, trees located over mostly impervious surface ( $\geq 50\%$ ) are classified as street trees. A local leaf area density (inside individual trees)  $L_{D,loc}$  of  $1 \text{ m}^3 \text{ m}^{-2}$  was used according to measurements over urban trees during summertime in Gothenburg, Sweden (Klingberg et al., 2017). The tree species measured by Klingberg et al. (2017) are assumed to be rather representative as they correspond to 6 of the 12 most common species present in Basel (Basel-Stadt, 2019). The density of street trees above the building column ( $L_{D,bld}$ , representing trees overhanging roofs) was assumed to be 0. This assumption was made giving that no reliable information was available about the overlapping between tree crowns and roofs. Further details regarding the methodology used to derive the urban canopy parameters are given in the Appendix A. The resulting  $L_{can}$  (Figure 4e) have values ranging from 0 up to  $3 \text{ m}^2 \text{ m}^{-2}$ , which is within the same range as reported in previous urban studies (Liss et al., 2010; Alonzo et al., 2015; Klingberg et al., 2017).

The thermal and physical properties of urban elements are listed in Table 2. The values are based on the recommendations of Loridan and Grimmond (2012) and are used as default values in COSMO-BEP-Tree. The values have already been evaluated

**Table 2.** Material properties used in the urban canopy model.

|   | Roof                       | Wall                       | Road                       |
|---|----------------------------|----------------------------|----------------------------|
| Albedo (-)  | <del>0.1</del> <u>0.10</u> | <del>0.1</del> <u>0.10</u> | 0.15                       |
| Emissivity (-)  | 0.85                       | <del>0.9</del> <u>0.90</u> | 0.95                       |
| Heat Capacity ( $\text{J m}^{-3} \text{K}^{-1} \times 10^6$ ) | <del>1.5</del> <u>1.50</u> | <del>1.4</del> <u>1.40</u> | <del>1.5</del> <u>1.50</u> |
| Heat Conductivity ( $\text{W m}^{-1} \text{K}^{-1}$ )         | <del>0.8</del> <u>0.80</u> | <del>1.0</del> <u>1.00</u> | <del>0.8</del> <u>0.80</u> |
| Total thickness (m)   | <del>0.5</del> <u>0.50</u> | <del>0.3</del> <u>0.30</u> | <del>1.0</del> <u>1.00</u> |
| <u>Roughness length (m)</u>                                   | <u>0.01</u>                | <u>~</u>                   | <u>0.01</u>                |

over Basel by Loridan and Grimmond (2012). A foliage scattering coefficient of 0.5 and a foliage emissivity of 0.95 are used according to Krayenhoff (2014).

A step function is used to model the stomatal conductance ( $g_s$ ) with values of  $140 \text{ mmol m}^{-2} \text{ s}^{-1}$  and  $20 \text{ mmol m}^{-2} \text{ s}^{-1}$  for daytime and night-time, respectively. The values are calculated combining the dataset of urban tree species in the public area of Basel (Basel-Stadt, 2019) with observed values of stomatal conductance (Konarska et al., 2016; Keel et al., 2007; Campbell and Norman, 2012; Xiong et al., 2018). Additional details are provided in the Supplement.

A uniform constant value of clumping index ( $\Omega$ ) of 0.5 was used for the reference simulation (STD). In the absence of easily applicable methods to determine  $\Omega$  in street canyon environments, we made our choice in analogy to measurements over an open Savannah (Ryu et al., 2010). Savannah is a natural environment that resemble the urban environment in terms of heterogeneity and openness of the tree canopy.

Anthropogenic heat flux ( $Q_A$ ) was estimated off-line with the Large scale Urban Consumption of energy model (LUCY; Allen et al. (2011); Lindberg et al. (2013)) version 2014a. LUCY simulates the hourly heat emissions from buildings, vehicles and human metabolism using gridded inventories of population density, primary energy consumption, air temperature, number of vehicles ~~and~~ seasonal, weekly, and diurnal cycles of traffic density. Gridded population density from Center for International Earth Science Information Network - CIESIN - Columbia University, International Food Policy Research Institute - IFPRI, The World Bank, and Centro Internacional de Agricultura Tropical - CIAT. 2011 (2011) with a resolution of 30" were used. ~~Monthly values~~ Values of energy consumption in the region of Basel for the months June and July 2015 have been obtained ~~for 2015 from Basel-Stadt~~ Basel-Stadt, replacing the default yearly average value, as recommended by Lindberg et al. (2013). The model estimated period-average values of  $Q_A$  of 4 and  $14.4 \text{ W m}^{-2}$  over the entire urban area and the city centre, respectively. A maximum value of  $22.9 \text{ W m}^{-2}$  was found around mid-day in the city centre. The values are consistent with previous estimates of  $20 \text{ W m}^{-2}$  as annual average from Christen (2005).

### 3.3 In situ measurements

Measurements from one 38 m tall flux tower and ~~six~~ five surface stations were used to evaluate the model results. Table 3 shows a summary of the corresponding observations. The site of Basel Klingelbergstrasse (BKLI) is an urban flux tower operated by

**Table 3.** Overview of observation sites used for model evaluation. Figure 3b shows the location of the sites.  $T$  is air temperature,  $q$  specific humidity,  $u$  wind speed,  $Q_H$  sensible heat flux, and  $Q_L$   $Q_E$  latent heat flux.

| Name | Variable <sup>a</sup>    | Type       | Height   | $f_{urb}$ | $L_{can}$ | LCZ                               |
|------|--------------------------|------------|----------|-----------|-----------|-----------------------------------|
| BKLI | $T, q, u, Q_H, Q_L, Q_E$ | Flux tower | 38 m     | 0.79      | 1.50      | Compact mid-rise (2)              |
| BFEL | $T$                      | Surface    | 2 m      | 0.98      | 1.03      | Compact mid-rise (2)              |
| BSJO | $T$                      | Surface    | 2 m      | 0.85      | 1.06      | Compact mid-rise <sup>b</sup> (2) |
| SLFR | $T$                      | Surface    | 2 m      | 0.51      | 0.35      | Large low-rise (8)                |
| BBIN | $T$                      | Surface    | 2 m      | <0.1 0.10 | 0.00      | Rural low plants <sup>c</sup> (D) |
| BLER | $T, q, Q_H, -Q_L$        | Surface    | 2 / 10 m | <0.1 0.10 | 0.00      | Rural low plants (D)              |

<sup>a</sup> Only variables used in this study are reported (the station may measure additional variables). <sup>b</sup> with scattered trees (LCZ B). <sup>c</sup> surrounded by open mid-rise residential areas (LCZ 2).

the ~~MRC-Lab~~ MCR-Lab (Meteorology, Climatology and Remote Sensing) of the University of Basel. BKLI is located in an area with a dense mix of mid-rise ~~building~~ buildings, corresponding to a Local Climate Zone (LCZ) 2 (Stewart and Oke, 2012), and relatively high density of street trees. The site consists of a 18 m high tower mounted on the flat roof of a 20 m high building. According to Lietzke and Vogt (2013) and Feigenwinter et al. (2012), the measurements can be considered to take

5 place above the roughness sub-layer, ~~but~~ and inside the inertial sub-layer. The footprint area for energy fluxes was estimated to ~~be~~ 0.60 km<sup>2</sup> for 90% of the source area in summer (Schmutz et al., 2016). The site comprises an eddy covariance system consisting of an ultrasonic anemometer (HS-50, Gill Instruments Ltd.) and an open-path CO<sub>2</sub>/H<sub>2</sub>O infra-red gas analyser (LI-7500, LI-COR Inc.). Air temperature is measured with a fine-wire thermocouple mounted in a fan-aspirated radiation shield (ASPTC, Campbell Scientific); relative humidity is measured with a mirror-type dew point hygrometer (THYGAN VTP6,

10 Meteorolabor AG); radiation fluxes are measured with a ~~net-four-component~~ radiometer (CNR4, Kipp & Zonen). It is worth to mention that the flux measurement is generally characterised by a considerable uncertainty (Richardson et al., 2012; Järvi et al., 2018). The uncertainty due to random errors is approx. 20% and 30% of the measured flux of sensible and latent heat, respectively (Hollinger and Richardson, 2005). Further details on the site and data processing are available in Schmutz et al. (2016); Lietzke and Vogt (2013); Feigenwinter et al. (2018).

15 Six surface stations are used to evaluate the model performance inside the urban canopy layer. The sites of Basel Feldbergstrasse (BFEL) and Basel St. Johannplatz (BSJO) are operated by the local environmental agency (Lufthygieneamt beider Basel). The site of BFEL is located in an area corresponding to compact mid-rise (LCZ 2) and has moderate density of street trees. It is ~~placed~~ positioned directly inside the street canyon at the intersection between two roads oriented along the 80°/260° and 160°/340° axis. Due to its vicinity to a south facing wall, BFEL may experience temperatures larger ~~then~~ than those ex-

20 pected in the ~~area~~ middle of the street canyon. The site of BSJO is located in a small park at the border of the old city core. The park is vegetated and extends to the north-east while the surrounding area is characterised by a dense mix of mid-rise buildings. Both sites BFEL and BSJO measure temperature with a thermometer with radiation shield (HC-S3, Campbell Scientific). The site of Saint-Louis Trois-Frontières (SLFR) is operated by the environmental agency of Alsace (ASP), France. The station is

located in the outskirts of the city in an area characterised by large low-rise buildings (LCZ 8). Few or no trees are present and the land cover is mostly paved or hard-packed. Air temperature is measured with a platinum resistance thermometer (HMP35D, Vaisala) with radiation shield. The site of Basel Binningen (BBIN) is operated by the Swiss Federal Office of Meteorology and Climatology (MeteoSwiss). The station is located in a rural/agricultural patch (LCZ D) at the south-western outskirts of the city. Low-rise residential areas surround the patch to the west, north and east. The surface underneath the sensor is natural (grass) and scattered low-rise buildings are present at 15-20 m distance. The site measures temperature with a thermometer with radiation shield and active ventilation (Rotronic HC2A-S3, Meteolabor). The site of Basel Lange Erlen (BLER) is operated by the ~~MRC-Lab as well~~ MCR-Lab. The station is located in a rural/agricultural patch (LCZ D) at 6 km distance from the city core. The surface underneath the sensor is natural (grass). Given the vegetated environment and distance from the city centre ~~and the vegetated environment~~, the site of BLER is used as a rural reference in this study. The station used the same instruments as the site BKLI for air temperature, relative humidity and radiation fluxes.

### 3.4 Satellite observations

Thermal infrared (TIR) data from the Landsat 7 are used to evaluate the modelled land surface temperatures ( $LST_{mod}$ ). Landsat acquires TIR data with a spatial resolution of 60 m. In contrast to other TIR satellite data (e.g. MODIS or SEVIRI), this resolution is fine enough to support high-resolution applications. Due to its small field of view, Landsat TIR retrievals are not affected by thermal angular anisotropy (Hu et al., 2016). During the investigation period, the Landsat 7 scene from 5 July 2015 was selected for the evaluation. No other cloud-free scenes from Landsat 7 or 8 were available during this period. Specifically, the Landsat 8 overpass of 4 July 2015 had to be discarded due to the presence of cirrus clouds.

The satellite surface temperature ( $LST_{sat}$ ) was derived from satellite data by correcting and converting the TIR signal received by the satellite sensor. The  $LST_{sat}$  calculation follows the approach described in Mussetti et al. (2020). We used the radiative transfer model MODTRAN (MODerate resolution atmospheric TRANsmission) to estimate the atmospheric transmission, the long-wave upwelling or atmospheric path radiance and the long-wave downwelling or sky radiance (Abreu and Anderson; Tardy et al., 2016). Atmospheric profiles from COSMO-BEP-Tree (STD simulation) were used in MODTRAN replacing its standard profiles. MODTRAN was used to transform the top-of-atmosphere radiance to surface-leaving radiance, which was further converted to LST by inverting the Planck function (Coll et al., 2010). A modified vegetation-threshold approach by Sobrino et al. (2008) was used to estimate the surface emissivity. As the original approach does not account for built-up areas, a lookup table for dark and bright urban surfaces with a range of 0.95 to 0.97, based on emissivity values for typical urban fabrics (Baldrige et al., 2009), was applied (Mitraka et al., 2012). ~~Using the above-described methods, the uncertainty in surface emissivity and atmospheric effects results in a total error of approximately 0.4 to 1.1 K (Jiménez-Muñoz and Sobrino, 2006).~~ Note that these values differ from those used in the model (Table 2). The uncertainty in the values of emissivity produces average changes in the LST of about 1.25 K (Supplement). On top of this, the LST is affected by uncertainties due to atmospheric correction, sensor and off-nadir view (Voogt and Oke, 2003; Chen et al., 2017).



The  $LST_{sat}$  has been re-mapped to the model grid of COSMO-BEP-Tree using an average re-sampling method, i.e. the average of all contributing pixels was computed (McInerney and Kempeneers, 2015). This re-sampling method ensures the conservation of the mean values.

$LST_{mod}$  was calculated combining the temperatures of roofs, streets, natural surfaces according to their respective fractions

5 as:

$$LST_{mod} = \sum_{d=1}^{N_d} T_{str} f_{dir} f_{str} (1 - CC) + \sum_{d=1}^{N_d} \overline{T_{veg}} f_{dir} f_{str} CC + \sum_{d=1}^{N_d} \left( \sum_{h=1}^{N_h} T_{roof} f_{roof} \right) f_{dir} f_{bld} + T_{nat} f_{nat}, \quad (14)$$

where  $T_{str}$  is the temperature of the streets (function of canyon direction ( $d$ )),  $f_{dir}$  is the fraction of canyons with direction  $d$ ,  $CC$  is the effective canopy cover from street trees,  $\overline{T_{veg}}$  is the vertically averaged temperature of the leaves of the street trees,  $T_{roof}$  is the temperature of the roofs (function of ( $d$ ) and height ( $h$ )) and  $T_{nat}$  is the temperature of the natural tile.  $CC$  is  
10 estimated from the leaf area density  $L$  using the relationship provided by Klingberg et al. (2017).

### 3.5 Performance metrics

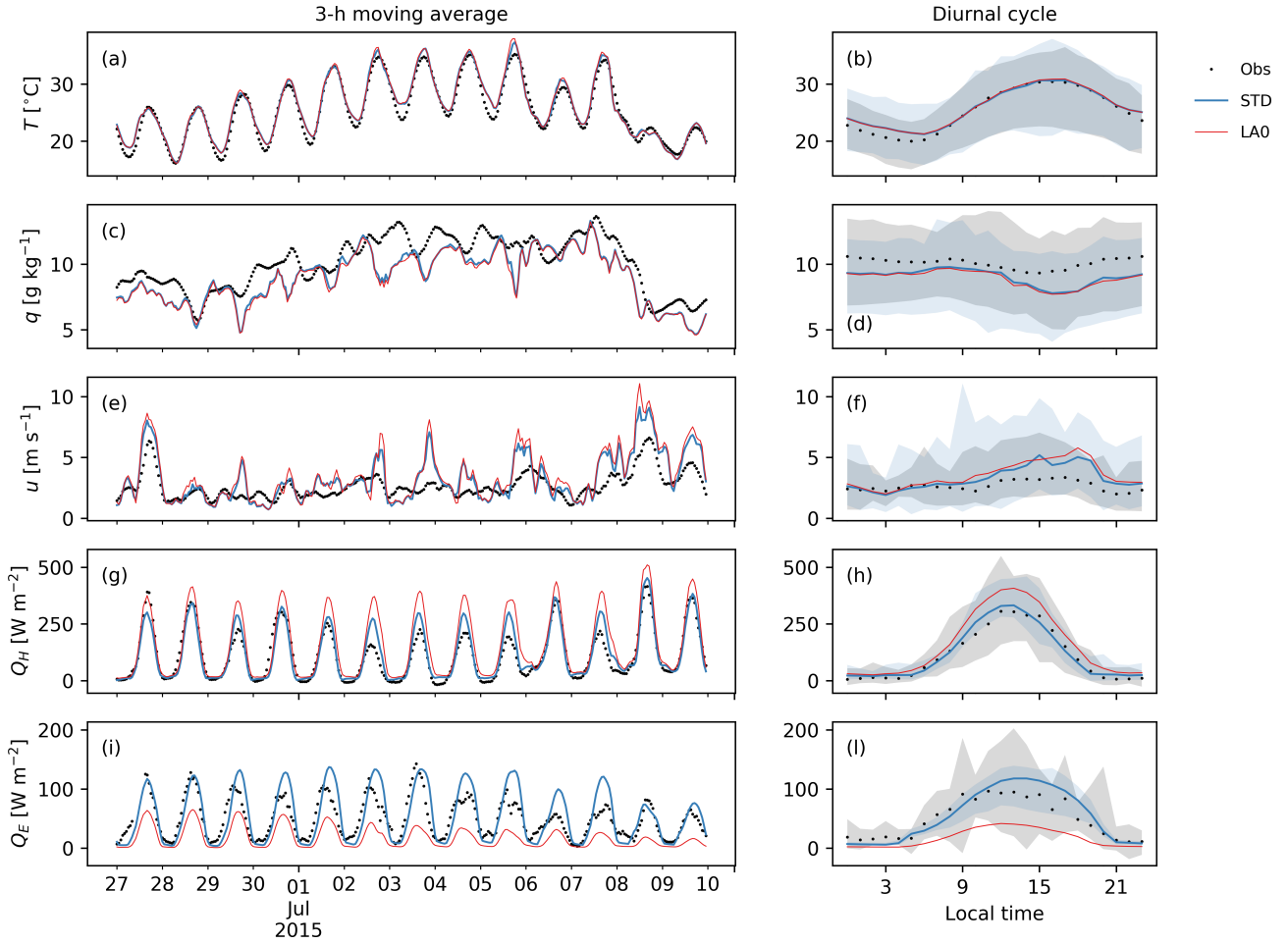
The model performance was evaluated against the observations using the following metrics: root-mean-square error (RMSE), systematic root-mean-square error ( $RMSE_s$ ), unsystematic root-mean-square error ( $RMSE_u$ ), mean bias error (MBE) and coefficient of determination ( $r^2$ ). This choice of metrics follows the recommendations of Willmott et al. (1985) and ensures  
15 comparability with previous urban climate model evaluations (e.g., Schubert and Grossman-Clarke, 2014; Oleson et al., 2008; Grimmond et al., 2011). The definition of the indices is available in Willmott et al. (1985); Willmott (1981).

## 4 Results

### 4.1 Evaluation against measurements from an urban flux tower

The comparison between model simulations and observations above the urban canopy layer at the BKLI flux tower site is  
20 shown in Fig. 5. Only the results corresponding to the reference simulation (STD) and to the simulation without street trees (LA0) are shown (see Table 1 for further details on the simulation set-ups). The statistical scores are given in Table 4. If not specified, we ~~will~~ refer to the STD simulation in the text below.

The model simulates the evolution of air temperature very well during the evaluation period (Figure 5a-a-b) although a slight overestimation with a MBE of 0.51 K and a systematic RMSE of 0.51 K was found. The overestimation occurs mostly during  
25 night-time ~~, an issue that was observed also in other urban climate studies with COSMO (Schubert and Grossman-Clarke, 2014; Wouters et al. 2017) and other models (e.g. Lee et al., 2011). The~~ and it seems to be related to an underestimation of the bulk albedo (Supplement) which can be traced back to the choice of material properties (Table 2). A too low albedo may have produced an excess in heat storage with consequently larger sensible heat release at night. Nevertheless, a night-time overestimation of ~~air temperature is attributed to the inability of COSMO to represent very stable boundary layer conditions (Cerenzia, 2017)~~ temperature was not  
30 found at the near-surface sites, indicating that more analyses are needed to better understand eventual issues with the choice of



**Figure 5.** Comparison between observations (Obs) and model simulations (STD, LA0) of air temperature (a,b), specific humidity (bc,d), wind speed (ee,f), sensible heat flux (dg,h) and latent heat flux (ei,l) at the site BKLI (38 m tall flux tower) during the selected simulation study period. The left (discarding the first 5 days as spin-up) and right (b,d,f,h,l) columns show the 3-h moving average and the diurnal cycle, respectively. Black dots indicate the observations. Blue and red lines indicate the results from the STD and LA0 model runs, respectively. The shaded areas represent the range of variability within the period for the observations and STD simulation only. Specifications of the model set-ups can be found in Section 3.1.

material properties. The impact of street trees on the air temperature above the canopy layer is found to be negligible, giving the very small differences between STD and LA0.

The evolution of specific humidity  $q$  throughout the simulation is generally well reproduced (Figure 5a-bc-d). The mean value of  $q$  is underestimated (MBE =  $-1.09 \text{ g kg}^{-1}$ ), which explains a large fraction of the systematic RMSE ( $\text{RMSE}_s = 1.13 \text{ g kg}^{-1}$ ). A similar pattern is found at the rural site of BLER (available in the Supplement), indicating that the bias is not related to the urban scheme. The bias would be consistent with a general underestimation of soil moisture in COSMO as reported by



**Table 4.** Root-mean-square errors (RMSE), mean-bias errors (MBE) and coefficient of determination ( $r^2$ ) of the simulated air temperatures ( $T_{\text{air}}$ ), specific humidity ( $q$ ), wind speed ( $u$ ), sensible heat flux ( $Q_H$ ) and latent heat flux ( $Q_E$ ) at the BKLI site (38 m above ground) during the study period 22 June 2015 – 9 July 2015 (discarding the first 5 days as spin-up). In addition to the total RMSEs (T), the systematic (S) and the unsystematic (U) RMSEs are listed. The statistics are provided for the standard simulation (STD) and the simulation without street trees (see Table 1 for details on the simulation set-up LA0). The units only apply to RMSE and MBE.

|                            |     | RMSE (T/S/U)  | MBE                     | $r^2$               |
|----------------------------|-----|---|-------------------------|---------------------|
| $T$ (K)                    | STD | 1.04 / 0.51 / <del>0.9</del> 0.90                                     | 0.51                    | 0.97                |
|                            | LA0 | <del>1.1</del> 1.10 / 0.58 / 0.93                                     | 0.58                    | 0.97                |
| $q$ (g kg <sup>-1</sup> )  | STD | 1.52 / 1.13 / 1.02  | -1.09                   | 0.74                |
|                            | LA0 | 1.61 / 1.24 / 1.03  | <del>-1.2</del> -1.20   | 0.73                |
| $u$ (m s <sup>-1</sup> )   | STD | 1.31 / <del>0.8</del> 0.80 / 1.03                                     | 0.61                    | <del>0.7</del> 0.70 |
|                            | LA0 | 1.62 / <del>1.1</del> 1.10 / 1.19                                     | 0.87                    | 0.69                |
| $Q_H$ (W m <sup>-2</sup> ) | STD | <del>40.1</del> 40.10 / 7.72 / <del>39.3</del> 39.30                  | <del>7.7</del> 7.70     | 0.88                |
|                            | LA0 | <del>74.8</del> 74.80 / <del>50.6</del> 50.60 / <del>55.2</del> 55.20 | <del>47.0</del> 47.00   | 0.84                |
| $Q_E$ (W m <sup>-2</sup> ) | STD | <del>25.1</del> 25.10 / 6.86 / <del>24.1</del> 24.10                  | 5.31                    | 0.70                |
|                            | LA0 | <del>37.8</del> 37.80 / <del>36.9</del> 36.90 / 8.57                  | <del>-32.2</del> -32.20 | 0.74                |

Davin et al. (2011). Another possible source of the bias may be related to the boundary conditions, with insufficient humidity advected from the boundaries of the domain. Street trees produce a very small increase in relative-specific humidity bringing the model in slightly closer agreement with the observations. This is quantified by a reduction in the systematic component of RMSE and MBE compared to LA0 of about 0.1 g kg<sup>-1</sup> (about 6 and 8 % of the entire RMSE and MBE, respectively).

- 5 The model simulates well the evolution of wind speed ( $u$ ) during the evaluation period (Figure 5e-de-f), but overestimates  $u$  during a few-several days and especially in the afternoon. The origin-of-the-overestimation-is-unclear-It-may-be-related-to-either-overestimation-was-not-limited-to-the-urban-area-but-was-also-found-at-the-rural-site (Supplement). Although its origin is unclear, possible causes may a misrepresentation of the synoptic-regional scale wind from the driving boundary conditions and an overestimation of the thermally driven flow induced by the surface temperature gradient between the city and the
- 10 surroundings. Nevertheless, the comparison with observations shows a small bias (MBE = 0.61 m s<sup>-1</sup>), lower than previous urban studies with COSMO (e.g. Schubert and Grossman-Clarke, 2014), and the unsystematic error dominates. Street trees reduce the simulated  $u$  and slightly improve the agreement with the observations. RMSE and MBE are reduced compared to LA0 of-by 0.31 and 0.26 m s<sup>-1</sup>, respectively.

- The sensible heat flux ( $Q_H$ ) follows very well the observations during the evaluation period (Figure 5e-fg-h) showing almost
- 15 no bias (MBE = 7.7 W m<sup>-2</sup>) but a RMSE of 40.1 W m<sup>-2</sup> with a large unsystematic component. These errors are comparable to, or smaller than those obtained in a recent inter-comparison of UCMs (Grimmond et al., 2011). The representation of street trees substantially improves the agreement with the observations. Street trees reduce simulated  $Q_H$  by as much as 100 W m<sup>-2</sup>

during mid-day. The improvements in RMSE and MBE compared to LA0 were as large as 34.7 and 39.3 W m<sup>-2</sup>, respectively. Additionally, an improvement of 0.06 in  $r^2$  was also found.

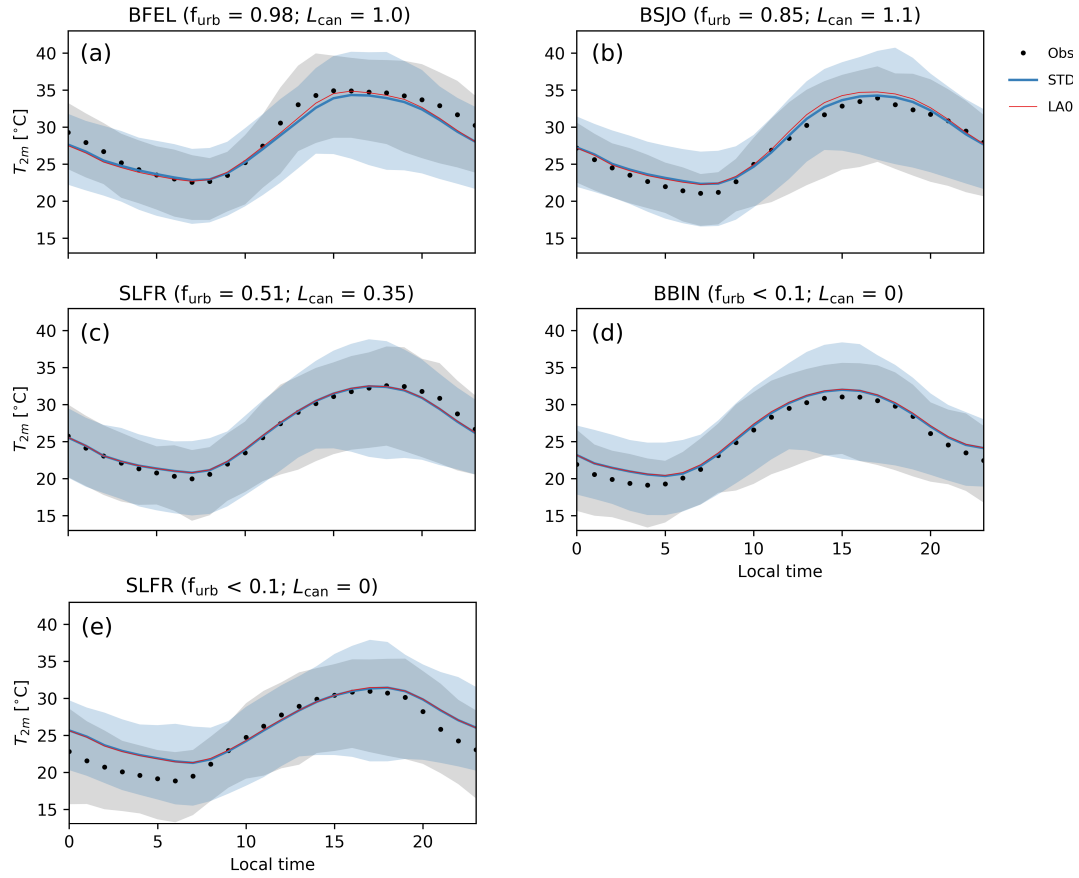
The evolution of latent heat flux ( $Q_E$ ) is generally well reproduced by the model (Figure 5g-hi-l), but the agreement is better during the initial and final phase of the heatwave when the synoptic forcing was stronger and wind speeds correspondingly higher. During the central part of the evaluation period, ~~a slight an~~ overestimation is found. Overall, the model showed a very small MBE of about 5 W m<sup>-2</sup> and a RMSE of 25.1 W m<sup>-2</sup>. Again, the scores are comparable to, or better than those obtained in a recent inter-comparison of UCMs (Grimmond et al., 2011). The slight overestimation during the central part of the simulation period could be caused by a limitation in soil water availability that may have been present in reality but was not considered in the simulations. Street trees had a very large impact on  $Q_E$  and improved the agreement with the observations substantially. RMSE and MBE improved compared to LA0 by 12.7 and 26.9 W m<sup>-2</sup>, respectively.

For completeness, the comparison between modelled and observed radiation fluxes is shown in the Supplement. The model simulations agree reasonably well with the observed radiation fluxes, although the upward short-wave radiation is underestimated. This indicates that the albedo is underestimated, which can be related to the choice of material properties (Table 2). The upward long-wave radiation is slightly overestimated, which can therefore be related to the underestimation of the albedo and, consequently, to an overestimation of surface temperatures. Overall, the representation of street trees had a very minor impact on the radiation fluxes.

## 4.2 Evaluation against surface stations

Within the urban canopy layer, the model was evaluated in terms of the 2-m air temperature ( $T_{2m}$ ) and canopy layer urban heat island (UHI) intensity. The UHI intensity is defined as  $UHI_i = T_i - T_{rur}$  where  $T_i$  is the 2-m air temperature at the urban site  $i$  and  $T_{rur}$  is the 2-m temperature at the rural reference site BLER. The comparison between model simulations and observations in term of ~~2-m air temperature~~  $T_{2m}$  is shown in Fig. 6. Only the results corresponding to the reference simulation (STD) and to the simulation without street trees (LA0) are shown. If not specified, we ~~will~~ refer to the STD simulation in the text below.

The model simulation compares generally well with the observed  $T_{2m}$ , especially in the urban area. The model errors are in the range of other urban climate studies (e.g. Schubert and Grossman-Clarke, 2014; Wouters et al., 2016). The difference between model and observations is more pronounced at the site BFEL where the systematic part of the RMSE is substantially higher ( $RMSE_s = 0.93$  K) than at other sites. This difference may be related to ~~the local~~ exposure of the ~~sensor to local influences~~ BFEL sensor, located in the vicinity of a (warm) south facing wall. Notable is also the overestimation of temperatures at the rural sites BLER and BBIN (MBE = 1.36 K and 0.88 K, respectively), which is most pronounced at night-time. This is ~~a well known issue of COSMO that occurs during very stable atmospheric conditions probably determined by the combined effect of underestimation in surface albedo and misrepresentation of atmospheric stability (Supplement).~~ A similar underestimation of albedo at rural sites has been already found by Schubert and Grossman-Clarke (2014). The underestimation is attributed to the default albedo scheme of COSMO-CLM, which fails to represent different vegetation types. It is unclear, however, whether this bias is due to specific conditions at the BLER site (grassland) or due to a general misrepresentation in the land-surface model of COSMO-CLM. The misrepresentation of stable boundary layer conditions



**Figure 6.** Comparison between period averaged daily profiles of observed (Obs) and model-simulated 2-m air temperature at the sites BFEL, BSJO, SLFR and BBIN and BLER during the selected study period (discarding the first 5 days as spin-up). Black dots indicate the observations. Blue lines and red lines indicate the results from the STD and LA0 model runs, respectively. The range of variability within the selected period is shown as shaded area for Obs (grey) and STD (light blue).

is a well-known issue of COSMO (Buzzi et al., 2011; Cerenzia, 2017) and has already been reported in previous studies (e.g. Schubert and Grossman-Clarke, 2014). (Schubert and Grossman-Clarke, 2014; Mussetti et al., 2020).

The representation of street trees produces a slight decrease in daytime air temperature, particularly around noon (Figure 6a-b). This effect is more evident at sites with higher urban fraction ( $f_{urb}$ ) and leaf are area index from street trees ( $L_{can}$ ).

5 The representation of street trees slightly reduced the RMSE and MBE at all the sites (but BFEL) compared to LA0. When averaged over all sites, the reductions in RMSE and MBE were 0.03 and 0.11 K, respectively.

The comparison between model simulation and observations in terms of UHI intensity for the simulations STD and LA0 is shown in Fig. 7. The statistical scores are presented in Table 6 for the simulation STD only. The model is able to capture the general features of the UHI, such as the daily evolution and the different magnitude of the UHI at the different sites, whereas the

10 temporal variability throughout the period is not always represented. Generally, the UHI intensity is underestimated, especially

**Table 5.** Root-mean-square errors (RMSE), mean-bias errors (MBE) and coefficient of determination ( $r^2$ ) of the simulated 2-m air temperature  $T_{2m}$  at the sites BFEL, BSJO, SLFR, BBIN and BLER during the [study period](#) ~~22 June 2015–9 July 2015 (discarding the first 5 days as spin-up)~~. In addition to the total RMSEs (T), the systematic (S) and the unsystematic (U) RMSEs are listed. The statistics are provided for the standard simulation (STD) and the simulation without street trees (LA0).

|                  |     | 2-m air temperature (K)                  |                              |       |
|------------------|-----|--|------------------------------|-------|
|                  |     | RMSE (T/S/U)                             | MBE                          | $r^2$ |
| BFEL             | STD | 1.66 / 0.93 / 1.37                       | -0.85                        | 0.94  |
|                  | LA0 | 1.59 / 0.72 / 1.41                       | <del>-0.7</del> <u>-0.70</u> | 0.94  |
| BSJO             | STD | 1.39 / 0.56 / 1.27                       | 0.56                         | 0.95  |
|                  | LA0 | 1.55 / 0.81 / 1.32                       | 0.77                         | 0.95  |
| SLFR             | STD | 1.03 / 0.19 / 1.01                       | 0.04                         | 0.97  |
|                  | LA0 | 1.03 / 0.18 / 1.02                       | 0.08                         | 0.97  |
| BBIN             | STD | 1.39 / 0.90 / 1.07                       | 0.88                         | 0.96  |
|                  | LA0 | 1.43 / 0.99 / 1.04                       | 0.98                         | 0.97  |
| BLER             | STD | 2.17 / 1.41 / 1.65                       | 1.36                         | 0.91  |
|                  | LA0 | <del>2.2</del> <u>2.20</u> / 1.45 / 1.65 | <del>1.4</del> <u>1.40</u>   | 0.91  |
| All <sup>a</sup> | STD | 1.53 / 0.80 / 1.27                       | 0.40                         | 0.95  |
|                  | LA0 | 1.56 / 0.83 / 1.29                       | 0.51                         | 0.95  |

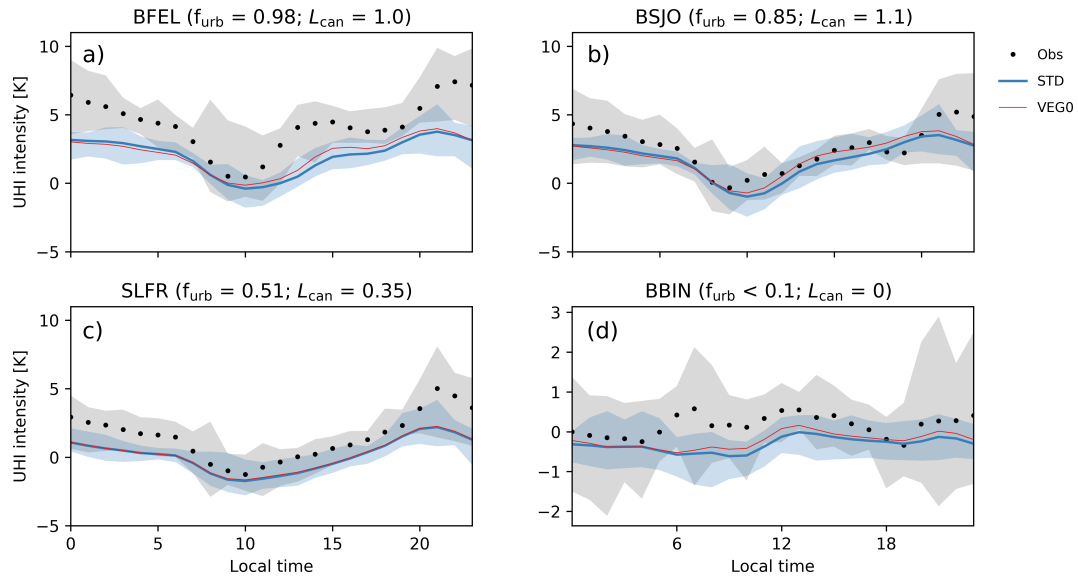
<sup>a</sup> average over all sites.

during night-time. This is related to the overestimation of night-time air temperatures at the rural site BLER (MBE = 1.36 K for STD) and, to a ~~lower~~minor extent, to the underestimation at the most urban sites (Table 5 and Supplement). The performance with respect to UHI intensity is better or comparable to similar studies (Wouters et al., 2016; Schubert and Grossman-Clarke, 2014).

- 5
- 10
- The representation of street trees<sup>’</sup> affects the simulated UHI intensity by producing a slight reduction in UHI intensity from late morning to early evening. The impact of street trees is more evident at sites with higher urban fraction and density of street trees. ~~The representation of street trees did not improve the model performance.~~ An improvement in model performance with street trees was, however, not evident in terms of UHI representation, ~~given that the warm bias at the rural site dominates.~~ Considering that street trees produce changes in the expected direction, this suggests that probably other model biases are masking the benefit of including street trees.

### 4.3 Evaluation against satellite observations of land surface temperature

Figure 8 compares the spatial distribution of simulated and remotely-sensed land surface temperature (LST) on 5 July 2015 at approx. 11:10 LT (10:10 UTC). Statistical scores are shown in Table 7.



**Figure 7.** Comparison between period-averaged daily profiles of observed (Obs) and model-simulated 2-m urban heat island (UHI) intensity at the sites BFEL, BSJO, SLFR and BBIN during the evaluation period. The site BLER is used as rural reference. Black dots indicate the observations. Blue and red lines indicate the results from the STD and LA0 model runs, respectively. The range of variability within the selected period is shown as shaded area (only for Obs and STD).

Overall, a good agreement between simulations and observations was obtained in terms of the domain averaged model bias and errors. As shown in Fig. 8, the spatial pattern of LST is generally well captured by the model indicating that the urban and rural land use variations are well represented. However, the model is found to underestimate the maximum LSTs and to overestimate the minimum LSTs at the outskirt-outskirts of the urban area (Figure 8d-e). Quantitatively, the model was found to have a small negative bias over the urban area (MBE = -0.39 K for STD) and a small positive bias when considering the entire domain (MBE = 0.34 K for STD). The RMSE was found to be lower in the urban area (RMBS = 2.71 K for STD) than in the entire domain. However, a relatively large component of RMSE was found to be systematic. The coefficient of determination was found to be only-quite low ( $r^2 = 0.58$  for STD). Only minor differences between STD and LA0 are seen, with STD being slightly colder over the urban area.

The moderate  $r^2$  and the relatively high component of systematic RMSE over urban areas may be explained by the fact that the same material properties (albedo, emissivity, heat capacity and heat conductivity) were used for the whole city. Therefore, spatial variations in material properties, e.g. between industrial and residential buildings, were not represented. This likely limits the model's ability to represent intra-urban gradients of surface temperature and, consequently, the related components such as sensible heat fluxes.

The explicit representation of street trees did not improve the agreement with satellite observations of land surface temperature. However, this observation should be taken with care given the uncertainty in the satellite observations of the order of 1 K. that satellite-derived LST in urban areas are known to have a large degree of uncertainty due to potential errors in the surface

**Table 6.** Root-mean-square errors (RMSE), mean-bias errors (MBE) and coefficient of determination ( $r^2$ ) of the simulated 2-m urban heat island (UHI) intensity (K) at the sites of BFEL, BSJO, SLFR and BBIN during the [study period](#)~~22 June 2015 – 9 July 2015 (discarding the first 5 days as spin-up)~~. The site of BLER is used as rural reference. In addition to the total RMSEs (T), the systematic (S) and the unsystematic (U) RMSEs are listed.

| 2-m UHI intensity (K) |     |                    |       |       |
|-----------------------|-----|--------------------|-------|-------|
|                       |     | RMSE (T/S/U)       | MBE   | $r^2$ |
| BFEL                  | STD | 2.7 / 2.53 / 0.95  | -2.26 | 0.57  |
|                       | LA0 | 2.6 / 2.43 / 0.92  | -2.12 | 0.56  |
| BSJO                  | STD | 1.56 / 1.2 / 1.0   | -0.86 | 0.54  |
|                       | LA0 | 1.53 / 1.12 / 1.04 | -0.65 | 0.48  |
| SLFR                  | STD | 1.8 / 1.63 / 0.75  | -1.37 | 0.67  |
|                       | LA0 | 1.77 / 1.6 / 0.75  | -1.33 | 0.67  |
| BBIN                  | STD | 0.93 / 0.86 / 0.36 | -0.48 | 0.43  |
|                       | LA0 | 0.86 / 0.78 / 0.37 | -0.38 | 0.45  |
| All <sup>1</sup>      | STD | 1.66 / 1.42 / 0.82 | -1.18 | 0.55  |
|                       | LA0 | 1.60 / 1.38 / 0.81 | -1.10 | 0.54  |

<sup>1</sup> average over all sites.

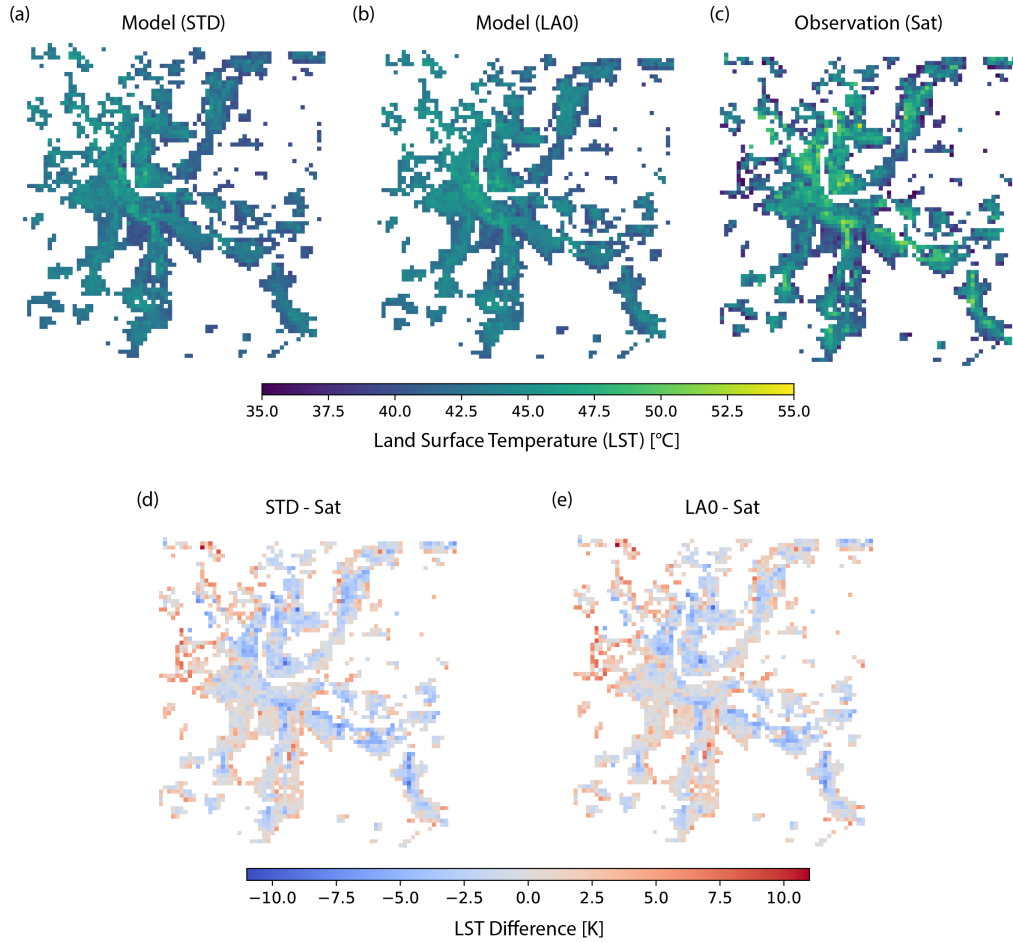
**Table 7.** Root-mean-square errors (RMSE), mean-bias errors (MBE) and coefficient of determination ( $r^2$ ) of the simulated (STD) land surface temperature (LST) [K] against the Landsat 7 satellite observations on the 5 July 2015 at approx. 11:10 LT (10:10 UTC). In addition to the total RMSEs (T), the systematic (S) and the unsystematic (U) RMSEs are listed. "Urb" represents the comparison only at the urban grid cells ( $f_{\text{urb}} > 0.1$ ). Grid cells contaminated by water or clouds have been removed.

| Surface Temperature (K) |     |  |       |                            |  |
|-------------------------|-----|--|-------|----------------------------|--|
|                         |     | RMSE (T/S/U)                             | MBE   | $r^2$                      |  |
| All (N=5853)            | STD | 3.14 / 2.36 / 2.07                       | 0.34  | <del>0.6</del> <u>0.60</u> |  |
|                         | LA0 | 3.08 / 2.25 / <del>2.1</del> <u>2.10</u> | 0.53  | 0.63                       |  |
| Urb (N=1930)            | STD | 2.71 / 2.36 / 1.34                       | -0.39 | 0.58                       |  |
|                         | LA0 | 2.61 / 2.25 / 1.31                       | -0.03 | 0.62                       |  |

[emissivity, sensor and satellite view bias \(Voogt and Oke, 2003; Chen et al., 2017\). Slight off-nadir view angles can introduce a bias over areas with tall buildings while LST variations at scales below 500 m may be unresolved \(Crawford et al., 2018\).](#)

Studies comparing simulated and remotely sensed surface temperatures over cities are rather scarce. We were only able to quantitatively compare our results with the studies of Wouters et al. (2016) and Hu et al. (2014). Other studies only performed

5 a qualitative comparison (Miao et al., 2009; Giannaros et al., 2013; Li and Bou-Zeid, 2013; Ramamurthy et al., 2017). The



**Figure 8.** Spatial distribution of land surface temperature (LST) [°C] on 5 July 2015 at approx. 11:10 LT (10:10 UTC). (a,b) show the simulated LST with STD and LA0 configurations, respectively; (c) shows the observed LST from Landsat 7 satellite; (d,e) show the LST differences between simulated and observed LST for the STD and LA0 configurations, respectively. Grid cells contaminated by the presence of the river or clouds have been removed.

MBE found in the present study is lower than what was found by Wouters et al. (2016) and Hu et al. (2014) for most of the conditions (different urban fractions and rural). Like the present study, Hu et al. (2014) found a general overestimation of daytime ST-surface temperature in their model while Wouters et al. (2016) reported an opposite behaviour. Wouters et al. (2016) and Hu et al. (2014) did not use other quantitative statistical metrics. Therefore, a comparison for RMSE and  $r^2$  was

5 not possible.

| (a)       | Daytime |       |       |       | (b)       | Night-time |       |       |       |
|-----------|---------|-------|-------|-------|-----------|------------|-------|-------|-------|
| $T_{air}$ | -0.51   | 0.55  | -0.22 | 0.25  | $T_{air}$ | -0.09      | -0.39 | -0.08 | 0.14  |
| $T_{str}$ | -0.58   | 0.56  | -0.04 | 0.04  | $T_{str}$ | 0.23       | -0.37 | -0.02 | 0.04  |
| $q$       | 0.51    | -0.53 | 0.26  | -0.33 | $q$       | 0.56       | -0.53 | 0.08  | -0.17 |
| $u$       | -0.29   | 0.32  | -0.02 | -0.01 | $u$       | -0.34      | 0.40  | -0.01 | -0.00 |
|           | LA+     | LA-   | SC+   | SC-   |           | LA+        | LA-   | SC+   | SC-   |

**Figure 9.** Model sensitivity  $S_{\phi}$  of 2-m air temperature  $T_{air}$ , street surface temperature  $T_{str}$ , specific humidity  $q$  and wind speed  $u$  to varying values of tree parameters (LA+, LA-, SC+, SC-). Values of  $S$  represent a spatial (over all urban grid cells) and temporal (over whole simulation period) average. Daytime corresponds to 10-18 local time, night-time to 21-5 local time.

#### 4.4 Model sensitivity to tree parameters

Figure 9 shows the model sensitivity of 2-m air temperature (called here  $T_{air}$ ), street surface temperature  $T_{str}$ , specific humidity  $q$  and wind speed  $u$  to varying values of tree parameters (LA+, LA-, SC+, SC-). The sensitivity  $S_{\phi,i}$  is defined here in relative terms normalised by the amplitude of the difference between a simulation with and without street trees:

$$S_{\phi,i} = \frac{\overline{\phi_i} - \overline{\phi_{STD}}}{|\overline{\phi_{LA0}} - \overline{\phi_{STD}}|}, \quad (15)$$

where  $\phi$  is a generic scalar field ( $T_{air}$ ,  $T_{str}$ ,  $q$ ,  $u$ ),  $\overline{\phi}$  represents the spatial and temporal average of the scalar field, and  $i$  represents the experiment as in Table 1. Generally, the model was more sensitive to changes in leaf area density (LA+, LA-) than to changes in stomatal conductance (SC+, SC-). During daytime (10-18 local time),  $T_{air}$  and  $T_{str}$  were the variables that were most sensitive to the choices of street tree parameters, whereas  $u$  was the least sensitive. At night (21-5 local time), conversely, the variable with the largest sensitivity was  $q$ .

An increase in the leaf area density from street trees (LA+) produced a reduction in daytime  $T_{air}$ ,  $T_{str}$  and  $u$ , and an increase in  $q$ . The reduction in  $T_{air}$  and  $T_{str}$  can be explained by the increase in interception of short-wave radiation from the denser foliage layer (see Eq. (1)) and, consequently, by the reduced amount of short-wave radiation reaching the wall and street surfaces within the canyon. The reduction in  $u$  can be explained by the augmented drag force exerted by the denser foliage layer (see Eq. (2)). The increase of  $q$  can be explained by the larger amount of leaf area available for transpiration (see Eq. (5)).

During night-time, the sensitivity experiment LA+ produced a marked increase in  $q$  and  $T_{str}$ .  $T_{air}$  remained approximately constant while  $u$  decreased. The responses of  $q$  and  $u$  follow the same mechanisms of the daytime response. The response of  $T_{str}$ , in contrast, is opposite to its daytime response. At night, enhanced interception of long-wave radiation emitted from



the internal surfaces of the canyon (walls and street) with increased leaf area density reduces the cooling of the street surface. Although enhanced tree foliage cools the street surfaces during the day due to shading, the heat capacity and corresponding thermal inertia of the street surfaces seems to be not sufficient to carry over this effect into the night-time to compensate for the reduced long-wave radiative cooling. Such an increase in night-time surface temperatures due to street trees, although somewhat counter-intuitive, has in fact been reported previously in observation-based studies (e.g. Shashua-Bar et al., 2009; Coutts et al., 2016).

A reduction in the leaf area density (LA-) generally produced an opposite response to that of an increase in leaf area density (LA+). An interesting exception is the reduction of  $T_{\text{air}}$  during-at night, suggesting that different processes with opposing sensitivities to values of LAD are at work. The largest night-time  $T_{\text{air}}$  reduction was achieved with LA-, which was associated with a similar reduction in  $T_{\text{str}}$ . A slight reduction in  $T_{\text{air}}$ , however, was also found for LA+, which contrasts with an increase in  $T_{\text{str}}$  for this scenario. This points towards complex, non-linear interactions between the different effects of LAD on radiation and winds. However, further studies are needed to corroborate this finding.

An increase in the stomatal conductance (SC+) produced an increase in daytime  $q$  and a decrease in  $T_{\text{air}}$ . A decrease (SC-) produced an opposite response. These effects are due to an increase in the source of moisture (see Eq. 5) and latent heat flux  $Q_E$  (see Eq. 4) as a response to the larger value of stomatal conductance  $g_s$ . The increase in  $Q_E$  causes a reduction in  $Q_H$  that consequently impacts  $T_{\text{air}}$ .  $T_{\text{str}}$  and  $u$  were found to be insensitive to changes in the stomatal conductance. This is expected since, in contrast to leaf area density, stomatal conductance has no influence on short-wave radiation or drag. During night-time, the model was found to be rather insensitive to changes in stomatal conductance. This is not surprising giving the low value of night-time stomatal conductance in the standard run (STD).

## 5 Future work

We acknowledge that there are a number of potential limitations in the current model system that need to be addressed in future studies. BEP-Tree, and therefore COSMO-BEP-Tree, employ a rather simplistic *stomata model* for representing the transpiration from street trees. In particular, constant and spatially homogeneous values of daytime and night-time stomatal conductance were employed ignoring the potentially important impact of water scarcity and the dependence on environmental conditions (e.g. air temperature and solar irradiation). In order to improve the representation of the transpiration from street trees and therefore the modelling of the associated latent heat fluxes, a mechanistic stomata model (e.g. Damour et al. (2010)) will have to be implemented. Furthermore, to properly represent soil water scarcity during extended drought periods, an urban hydrology module should be developed and coupled to COSMO-BEP-Tree (e.g. Järvi et al., 2011; Yang et al., 2015; Stavropoulos-Laffaille et al., 2018)

In this study, a homogeneous and constant value of the clumping index ( $\Omega$ ) was used. The value was chosen from a measurement over a natural environment that resembles the urban environment in terms of openness and heterogeneity of the tree canopy. In order to improve urban climate simulations with street trees, a methodology to estimate the spatial distribution of

$\Omega$  over urban environments needs to be established. A future implementation of the clumping effects should also consider the temporal changes of  $\Omega$  ~~that is known to vary~~ with the solar zenith angle (Ryu et al., 2016).

The *coupling approach* employed in this work to link COSMO-CLM and BEP-Tree is based on model integration (also called joint coupling by Brandmeyer and Karimi (2000)) rather than on an independent coupler. The model integration approach has the advantage of maximising the computational performance by reducing the information exchanged. As a drawback, it does not support independent model updates. In other words, every update in the standalone version of each individual model (COSMO-CLM and BEP-Tree) potentially requires additional effort in re-building the coupled version. In order to allow flexible coupling between future versions of COSMO-CLM and BEP-Tree, the use of a coupler would be interesting. As an example, COSMO-CLM has been already used with the OASIS coupler (Craig et al., 2017) to couple with different land surface models (Will et al., 2017).

Even if generally comparable with other models, the *performance* of COSMO-BEP-Tree still has room for improvements. The model evaluation against urban flux tower measurements revealed a systematic underestimation of specific humidity ( $q$ ). The source of this bias needs to be better investigated by, for instance, (a) analysing the soil moisture content and (b) evaluating  $q$  at the model boundaries provided by another model. The model evaluation ~~at 2-m~~ reveals an overestimation of ~~the air temperature~~ night-time air temperature above the urban canopy layer, which can be due to the use of default values of material properties. Additionally, the 2-m air temperature is overestimated at the rural sites especially at night. ~~This issue has been already found in other studies (e.g. Schubert and Grossman-Clarke, 2014) and is attributed to the inability of the model to reproduce very stable atmospheric conditions (Buzzi et al., 2011). More~~ Beside previous studies where this bias was attributed only to the misrepresentation of stable boundary layer conditions (Mussetti et al., 2020), the present study provides new evidences that point out to an underestimation of albedo over rural areas. Future studies need to address this issue that partially limits the model ability to represent the urban heat island effect. Potential solutions include the use of more advanced representations of albedo, already available as options in COSMO-CLM, with explicit consideration of vegetation type or satellite-based albedo values (Tölle et al., 2018). In respect to the representation of stable boundary layer conditions, more recent versions of the COSMO model (from version 5.4a onward) ~~include a~~ promise a better performance thanks to a revised turbulence scheme ~~that has the potential to better represent such conditions~~.

In the current study, spatially homogeneous values of *material properties* of roof, wall and street elements were used. Therefore, the impact of the spatial variation of material properties between areas with different building types (e.g. old town vs. industrial areas) could not be reproduced. Further versions of UCPgenerator need to consider different urban classes with distinct material properties. This could be achieved, for example, by making use of the data from the World Urban Database and Access Portal Tools (~~WUDAP~~ WUDAPT, Ching et al. (2018)). It has to be noted that the model resolution in this study falls within the so-called "grey zone" or "terra incognita" (Wyngaard, 2004), a condition where the horizontal grid spacing and hence the spatial filter used on the equations of motion are of the same order of the energy- and flux-containing turbulence. This may produce a misrepresentation in the turbulence and wind speed patterns (Efsthathiou and Beare, 2015). High-resolution urban climate studies reported the formation of horizontal convective rolls which have an unclear counterpart in the real atmosphere (Salamanca et al., 2012; Gutiérrez et al., 2015). Although such issues were not visible in this application, these

aspects need to be considered in future studies, for example by exploring the performance of new scale-aware turbulence models (Ito et al., 2015; Shin and Hong, 2015, e.g.) or large-eddy-simulation schemes as recently developed for the COSMO model (Panosetti et al., 2016)."

## 6 Conclusions

5 An urban climate model with explicit representation of street trees was developed featuring a two-way coupling between the vegetated urban canopy model BEP-Tree (Krayenhoff et al., 2019) and the mesoscale weather and climate model COSMO-CLM (COSMO, 2019). The coupled model, named COSMO-BEP-Tree, mechanistically represents the heat, momentum, turbulence and moisture interactions between street trees, urban elements and the atmosphere during realistic weather and climate simulations. COSMO-BEP-Tree comes with a tool to pre-process all data required by BEP-Tree in addition to the standard  
10 inputs for COSMO.

The model's performance and applicability were demonstrated for the urban area of Basel during a three week heatwave event in June-July 2015. The model evaluation made use of measurements from an urban flux tower, a network of surface stations, and land surface temperature observations from the Landsat 7 satellite. The application demonstrated that the model successfully captures essential features of the urban heat island during such a heat wave event and that city-wide interactions  
15 between street trees, urban elements and the atmosphere are represented in a realistic way.

In particular, the model captured the day-to-day as well as the mean diurnal evolution of air temperature, relative humidity, wind speed, sensible and latent heat fluxes relatively well at all observation sites. The comparison with measurements above the urban canopy layer revealed very small model biases and errors for air temperature, wind speed, sensible and latent heat fluxes. The model underestimated the specific humidity and slightly overestimated wind speed in the afternoon.

20 The comparison with measurements of air temperature within the urban canopy layer ( $T_{2m}$ ) revealed a generally good model performance. At rural sites, the model showed an overestimation of  $T_{2m}$  as already reported in previous urban climate studies using the COSMO model. The model was able to reproduce the diurnal cycle of the canopy layer urban heat island effect (UHI) generally well~~while~~, although its magnitude was underestimated during night-time. This issue was largely due to the warm night-time bias at the rural reference site. At the most urban site, ~~it~~this difference is additionally explained by an  
25 ~~underestimation-overestimation~~ of the night-time temperature.

In terms of the spatial pattern of the land surface temperature (LST), the model was found to have a relatively small RMSE and very small MBE compared to the satellite observations. However, the spatial variability over the urban area and especially some local peaks in LST were not fully captured by the model, which could be related to the use of spatially homogeneous material properties. Nevertheless, it should be noted that satellite LST in urban areas are associated with a large degree of  
30 uncertainty.

The inclusion of street trees in the model generally improved the performance against observations of  $Q_H$ ,  $u$ ,  $q$  and  $Q_E$ . In the case of  $T_{2m}$  this was found only ~~in some of the~~ at some evaluation sites. The representation of street trees had a large effect on sensible and latent heat fluxes and wind speed but only a minor influence on specific humidity and air temperature above

the urban canopy layer. Street trees were found to reduce  $Q_H$  and  $u$ , and to increase  $q$  and  $Q_E$ . Within the canopy layer, a slight decrease in  $T_{2m}$  during daytime was found when including street trees, and a very small increase in  $T_{2m}$  during night-time. This behaviour is in agreement with measurement studies over street canyons with trees (Shashua-Bar et al., 2009; Coutts et al., 2016). In terms of UHI intensity, this results in a slight reduction during daytime and a very small increase during night-time.

- 5     The model responded realistically to changes in the street ~~trees~~-tree parameters: leaf area density and stomatal conductance. During daytime, a 50% increase in leaf area density produced a reduction in air temperature, street temperature and wind, and an increase in humidity. At night, the same 50% increase in street ~~trees~~-tree density produced mainly an increase in humidity. A 50% increase in stomatal conductance produced an increase in humidity and a decrease air temperature.

- 10     This study demonstrates the ability of COSMO-BEP-Tree to simulate the impact of street trees for whole cities in a realistic way, by considering the multiple interactions between trees, urban elements and the atmosphere from the scale of individual street canyons to the scale of regional weather patterns. By including the effect of street trees, COSMO-BEP-Tree significantly enhances the modelling capabilities with respect to its predecessor COSMO-DCEP, adding an important dimension to the study of UHI mitigation and climate change adaptation strategies. Furthermore, the improved agreement with observations clearly underlines the importance of considering street trees in such simulations.

- 15     *Code and data availability.* COSMO-CLM is freely available for scientific usage to the members of the CLM-Community. The procedure to become a member is available on the community website (<https://www.clm-community.eu/>) or by contacting the coordination office ([clm-coordination@dwd.de](mailto:clm-coordination@dwd.de)). The code of COSMO-BEP-Tree is made freely available to any user with a valid CLM-Community membership (contact Dominik Brunner: [dominik.brunner@empa.ch](mailto:dominik.brunner@empa.ch)). All research data and scripts to produce figures and analysis are archived on Zenodo (Mussetti, 2019a). UCPgenerator v1.0 can be downloaded from the GitHub repository (Mussetti, 2019b). The LUCY model, used to calculate  
20     the anthropogenic heat flux, can be downloaded from Grimmond et al. (2018). BEP-Tree as a standalone model can be obtained by contacting Scott Krayenhoff ([skrayenh@gmail.com](mailto:skrayenh@gmail.com)). MODTRAN is a commercial software: <http://modtran.spectral.com/> (accessed 26-08-2019).

**Table A1.** Urban canopy parameters required as input to COSMO-BEP-Tree.

| Variable           | Description                              | Dimension          | NetCDF Name                   |
|--------------------|--|--------------------|-------------------------------|
| $f_{\text{urb}}$   | Planar fraction of urban surfaces        | lon, lat           | FR_URB                        |
| $W_{\text{can}}$   | Average canyon width [m]                 | lon, lat, dir      | W_STREET                      |
| $W_{\text{bld}}$   | Average building width [m]               | lon, lat, dir      | W_BUILD                       |
| $f_{\text{dir}}$   | Canyon direction distribution            | lon, lat, dir      | FR_UDIR                       |
| $\gamma$           | Probability to have a building           | lon, lat, dir, hgt | FR_ROOF                       |
| $L_D$              | Leaf area density in the canyon column   | lon, lat, dir, hgt | LAD_C                         |
| $L_{D,\text{bld}}$ | Leaf area density in the building column | lon, lat, dir, hgt | LAD_B                         |
| $\Omega$           | Clumping index                           | lon, lat           | OMEGA_R, OMEGA_D <sup>a</sup> |

<sup>a</sup> BEP-Tree gives the option to define a separate clumping coefficient for the drag and turbulence parameterisations.

## Appendix A: UCPgenerator v1.0

UCPgenerator v1.0 is a python program that generates the additional input fields required by COSMO-BEP-Tree (Mussetti, 2019b). The list of the additional input field required by COSMO-BEP-Tree is provided in Table A1. First of all, the urban fraction  $f_{\text{urb}}$ , defined as the planar fraction of urban surfaces in a model grid cell, is estimated from the given imperviousness dataset (see Section 2.4 for further details). A threshold method is used estimate the average urban fraction in the COSMO-BEP-Tree grid cell (function `urbfrac`). `urbfrac` performs also the reference system transformation.

The parameters that characterise the canyon geometry are then estimated in the function `cangeom` from the building dataset. Planar and frontal area density are calculated according to Grimmond and Oke (1999) as:

$$\lambda_p = \frac{\overline{A_P}}{\overline{A_T}} \quad (\text{A1})$$

10 and

$$\lambda_f = \frac{\overline{A_F}}{\overline{A_T}}, \quad (\text{A2})$$

where  $\lambda_p$  is the planar area density,  $\lambda_f$  is the frontal area density,  $\overline{A_P}$  is the plan area of building elements,  $\overline{A_T}$  is the total horizontal surface area and  $\overline{A_F}$  is the frontal area of building elements. Canyon and building widths are calculated according to Martilli (2009) as:

$$15 \quad W_{\text{bld}} = \frac{\lambda_p}{\lambda_f} h \quad (\text{A3})$$

and

$$W_{\text{can}} = \left( \frac{1}{\lambda_p} - 1 \right) \frac{\lambda_p}{\lambda_f}, \quad (\text{A4})$$

where  $W_{\text{can}}$  is the canyon width [m],  $W_{\text{bld}}$  is the building width [m] and  $h$  is the average building height [m].

Each building in the building dataset is clustered according to [height](#) and direction classes that can be set by the user through the namelist of COSMO-BEP-Tree. The vertical distribution of buildings  $\gamma$  and the distribution of canyon directions  $f_{\text{dir}}$  are determined from the above mentioned clustering.

The sub-routine `veget` finally computes the grid cell averaged values of  $L_{\text{D,can}}$  and  $L_{\text{D,nat}}$ . It is designed to read as  
5 input a LIDAR based dataset containing the height distribution of street trees. Making assumptions on the specific leaf area density inside an average tree ( $L_{\text{D,loc}}$ ), the routine calculates the total  $L_{\text{D}}$  in each model grid cell. If a dataset of all trees (e.g. including trees in parks) is available, the sub-routine identifies the street trees by using  $f_{\text{urb}}$ . Specifically, trees located over mostly impervious surface ( $f_{\text{urb}} > 0.5$ ) are classified as street trees (contributing to  $L_{\text{D,can}}$ ). The rest of the trees are classified as natural trees and contribute to  $L_{\text{D,nat}}$ . The routine also supports user defined modifications to the  $L_{\text{D}}$  to simulated greening  
10 or clearing scenarios.

UCPgenerator v1.0 requires the 3D building dataset to be in the format of ESRI shapefiles (.shp) with a level of detail (LoD) of 1. In LoD 1 datasets, buildings are represented as prismatic blocks without information on the roof structure (Kolbe et al., 2005).

*Author contributions.* GM conceived the work, performed the simulations and analysed the results. SH and GM implemented the coupling  
15 between COSMO and BEP-Tree. GM, DB, SH, ESK interpreted the results. ESK and SS provided the source codes of BEP-Tree and COSMO-DCEP, respectively. CF, RV and AW provided some of the input data and observations. DB, SH, JA and JC supervised the work. GM wrote the manuscript with contributions from all the co-authors.

*Competing interests.* The authors declare that they have no conflicts of interest.

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# Supplement to "COSMO-BEP-Tree v1.0: a coupled urban climate model with explicit representation of street trees"

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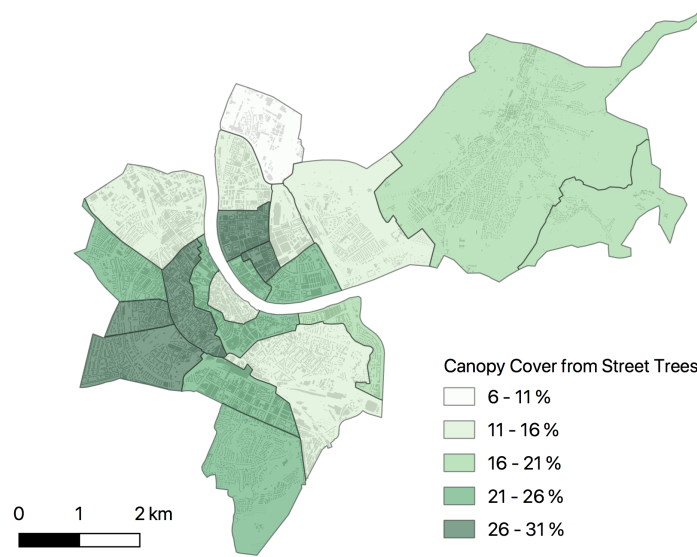
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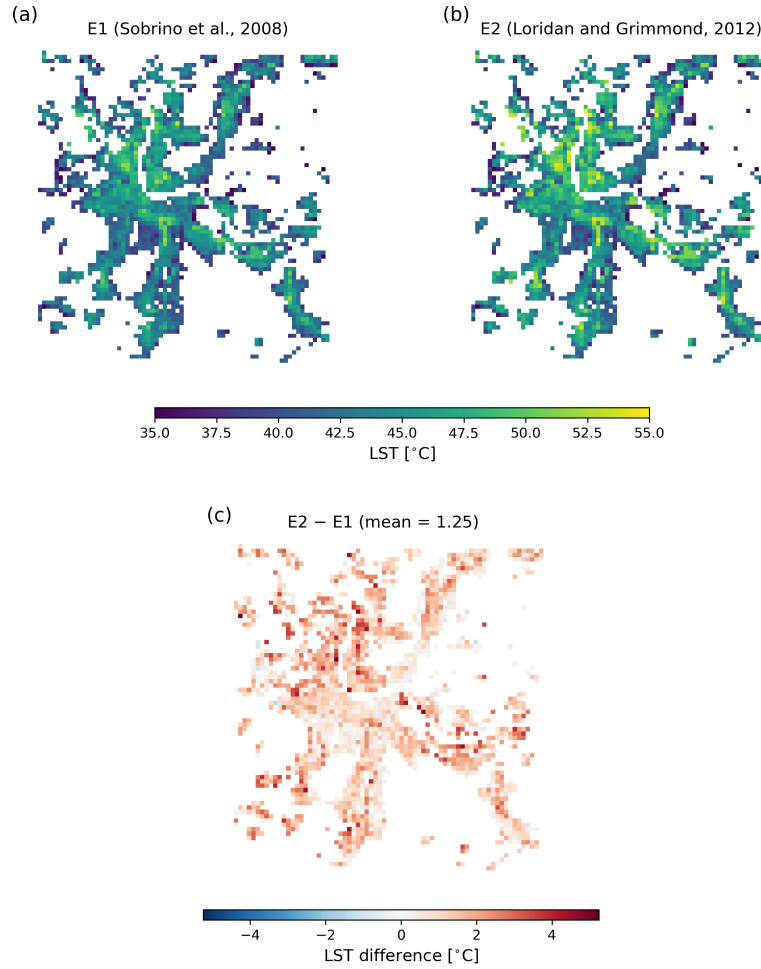


**Figure 1.** [Canopy cover from street trees in the administrative districts of Basel-Stadt \(Geoportal Kanton Basel-Stadt, 2019\).](#) [Building geometries are shown in the background \(Federal Office of Topography, 2007\).](#)

**Table 1.** Number and percentage of the 10 most common tree species found as public street tree in the inner part of Basel. The data are derived from the inventory of urban trees of the city of Basel (Geoportal Kanton Basel-Stadt, 2019).

| Species  | Num. | %    | $g_s$ [mmol m <sup>-2</sup> s <sup>-1</sup> ] |                  |
|--|------|------|---|------------------|
|  |      |      | night   | day              |
| <i>Tilia x euchlora</i> (Caucasian lime)       | 3433 | 11.3 | 20 <sup>a</sup>                               | 100 <sup>a</sup> |
| <i>Platanus x hispanica</i> (London plane)     | 2766 | 9.1  | 18 <sup>b</sup>                               | 180 <sup>b</sup> |
| <i>Aesculus hippocastanum</i> (Horse-chestnut) | 2083 | 6.9  | 20  | 90               |
| <i>Acer platanoides</i> (Norway maple)         | 1495 | 4.9  | 18  | 180              |
| <i>Tilia cordata</i> (Small-leaved lime)       | 990  | 3.3  | 20 <sup>a</sup>                               | 100 <sup>a</sup> |
| <i>Carpinus betulus</i> (Common hornbeam)      | 903  | 3.0  | 20 <sup>c</sup>                               | 100 <sup>c</sup> |
| <i>Acer campestre</i> (Field maple)            | 851  | 2.8  | 18 <sup>b</sup>                               | 180 <sup>b</sup> |
| <i>Pinus sylvestris</i> (Scots pine)           | 682  | 2.3  | 17 <sup>d</sup>                               | 330 <sup>d</sup> |
| <i>Quercus robur</i> (Common oak)              | 671  | 2.2  | 22  | 210              |
| <i>Taxus baccata</i> (Common yew)              | 651  | 2.1  | 15 <sup>e</sup>                               | 80 <sup>e</sup>  |
| Average <sup>f</sup>                           |      |      | 19  | 142              |

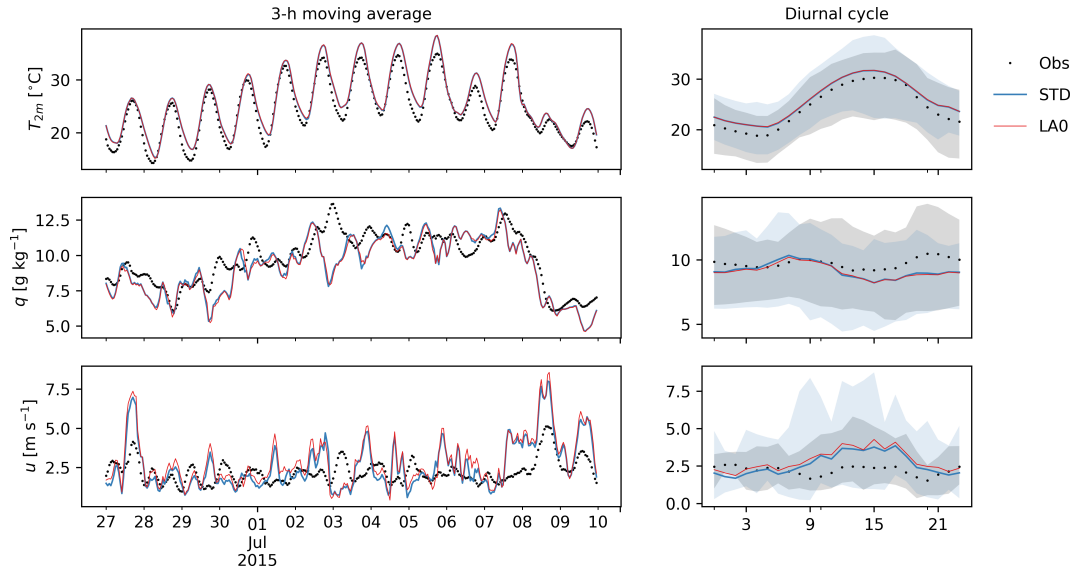
<sup>a</sup> values of *Tilia europaea*. <sup>b</sup> values of *Acer platanoides*. <sup>c</sup> from Keel et al. (2007) Fig. 2 year 2003. For night-time, the value from *Tilia europaea* are used. <sup>d</sup> from Campbell and Norman (2012) Table 7.2 values for *Pinus monticola*. <sup>e</sup> from Xiong et al. (2018) Fig. 2. <sup>f</sup> weighted with the corresponding % of trees within the sample.



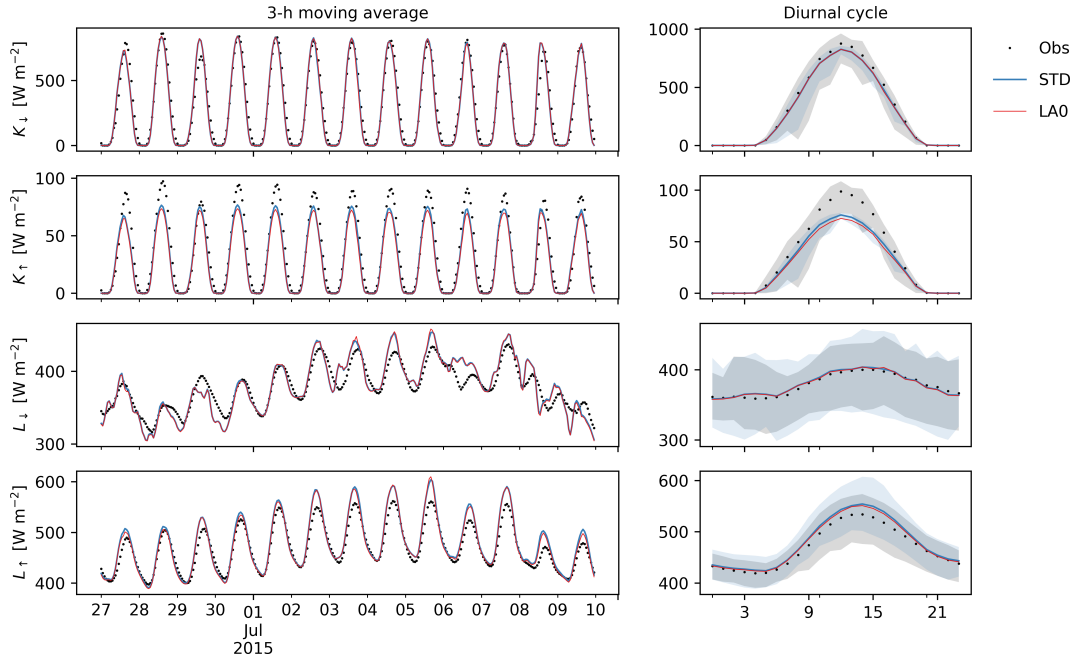
**Figure 2.** Impact of the choice of emissivity values for urban surfaces on the land surface temperature (LST). (a,b) show the observed LST with emissivity values from Sobrino et al. (2008) and Loridan and Grimmond (2012), respectively; (c) shows the difference in LST between the two approaches.

**Table 2.** Root-mean-square errors (RMSE), mean-bias errors (MBE) and coefficient of determination ( $r^2$ ) of the observed air temperatures ( $T_{air}$ ), specific humidity ( $q$ ), and wind speed ( $u$ ) at the rural site of BLER (10 m above ground) during the period 22 June 2015 - 9 July 2015 (discarding the first 5 days as spin-up). In addition to the total RMSEs (T), the systematic (S) and the unsystematic (U) RMSEs are listed. The statistics are provided for the standard simulation (STD) only.

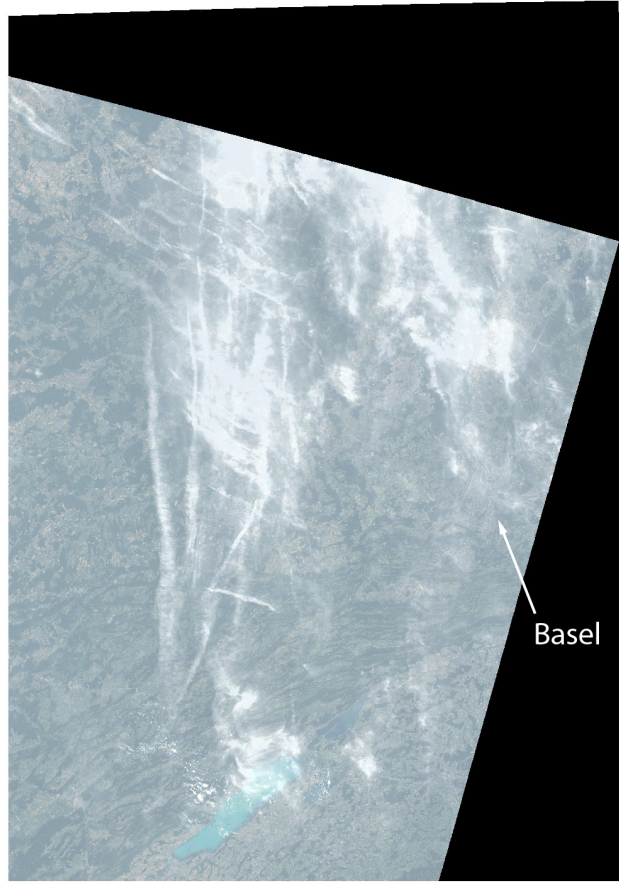
|                      | RMSE (T/S/U)       | MBE   | $r^2$ |
|----------------------|--------------------|-------|-------|
| $T_{air}$ (K)        | 1.74 / 1.34 / 1.10 | 1.34  | 0.96  |
| $q$ ( $g\ kg^{-1}$ ) | 1.33 / 0.60 / 1.19 | -0.56 | 0.65  |
| $u$ ( $m\ s^{-1}$ )  | 1.32 / 0.43 / 1.25 | 0.41  | 0.32  |



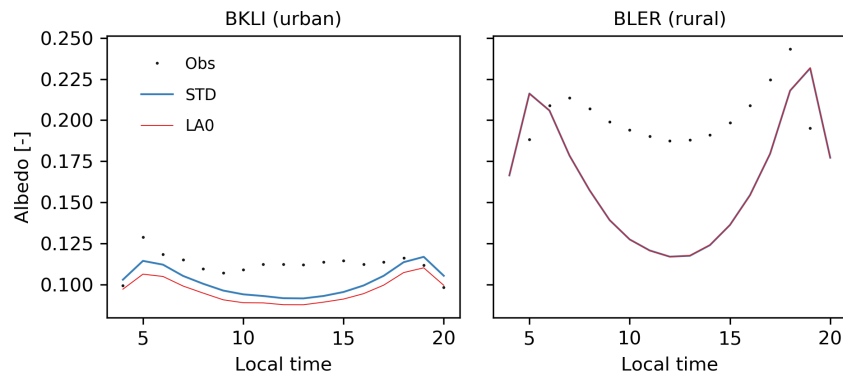
**Figure 3.** Comparison between observations (Obs) and model simulations of air temperature (a), specific humidity (b), wind speed (c), sensible heat flux (d) and latent heat flux (e) at the rural reference site BLER (10 m above ground) during the selected period (discarding the first 5 days as spin-up). Blue lines and red lines indicate the results from the STD and LAD0 model set-ups, respectively. The shaded areas represent the range of variability within the period for the observations and STD simulation only.



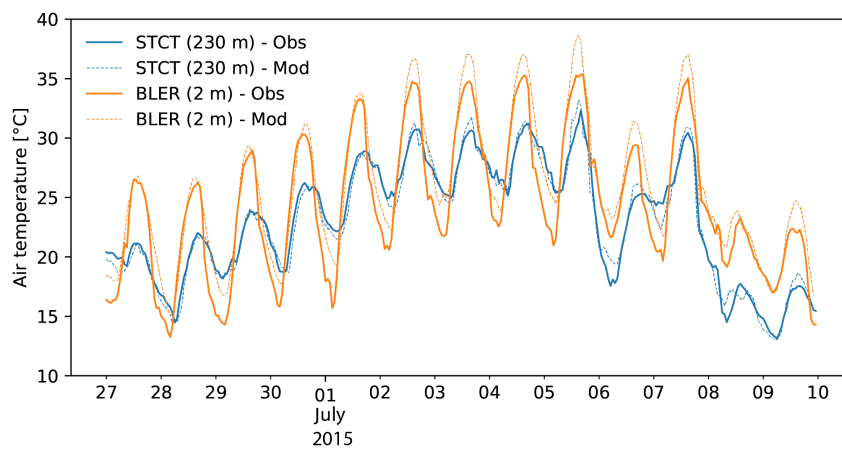
**Figure 4.** Comparison between observations (Obs) and model simulations of downward short-wave radiation  $K_{\downarrow}$  (a), upward short-wave radiation  $K_{\uparrow}$  (b), downward long-wave radiation  $L_{\downarrow}$  (c) and upward long-wave radiation  $L_{\uparrow}$  (d) at the site of BKLI during the selected period (discarding the first 5 days as spin-up). Blue lines and red lines indicate the results from the STD and LAD0 model set-ups, respectively. The shaded areas represent the range of variability within the period for the observations and STD simulation only.



**Figure 5.** Cirrus clouds above the model domain detected from the cirrus band (Band 9) of Landsat 8 satellite during the overpass of 4 July 2015.



**Figure 6.** Average diurnal profile of observed (Obs) and modelled (STD, LA0) albedo. Albedo is calculated as the ratio between upward and downward short-wave radiation. The site BKLI is characterised by an urban fraction ( $f_{urb}$ ) of 0.79.



**Figure 7.** Air temperature evolution at the sites of St. Chrischona (STCT, 230 m a.g., lat=47.571767, lon=7.687094) and BLER (2 m a.g.). Solid and dashed lines indicate observations (Obs) and model simulation (Mod), respectively. The evolution suggests the formation of a stable boundary layers (characterised by a strong temperature inversion) during the nights in central part of the period.

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