

1 **Supplement of**

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3 **ORCHIDEE MICT-LEAK (r5459), a global model for the production, transport and**
4 **transformation of dissolved organic carbon from Arctic permafrost regions, Part**
5 **2: Model evaluation over the Lena River basin.**

6
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19 **Text S1: Groundwater DOC Concentrations**

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21 The high groundwater reservoir DOC concentrations simulated in high altitude regions
22 by ORCHIDEE MICT-L is related to the fact that, in the model, DOC is rapidly produced
23 and infiltrated deep into soil above the permafrost table, to the point that it reaches the
24 simulated groundwater pool relatively quickly, allowing it to enter this reservoir before
25 being metabolised through the soil column –hence allowing for the relatively high
26 groundwater concentrations found in mountain areas. Because of the prevailing low
27 temperatures, this DOC is not quickly decomposed by microbes and instead feed the
28 groundwater DOC pool.

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31 **Text S2: Evaluation of Simulated NPP and Soil Respiration**

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33 Rates of yearly net primary production (NPP) for Russian and Siberian forests have been
34 inferred in situ from eddy flux and inventory techniques to range from 123-250 gC m⁻²
35 yr⁻¹ (Beer et al., 2006; Lloyd et al., 2002; Roser et al., 2002; Schulze et al., 1999;
36 Shvidenko and Nilsson, 2003). We likewise simulate a broad range of NPP carbon
37 uptake rates, of 61-469 gC m⁻² yr⁻¹ averaged per grid cell over the Lena basin, with a
38 mean value of 210 gC m⁻² yr⁻¹. NPP is heterogeneously distributed over space and
39 between PFTs (SI, Fig. S5c), with forests averaging 90 gC m⁻² yr⁻¹ and grasslands
40 averaging 104 gC m⁻² yr⁻¹ over the basin as a whole. Low values tended to originate in
41 basin grid cells with elevated topography or high mean slope, while the maximum value
42 was standalone, exceeding the next greatest by ~100 gC m⁻² yr⁻¹, and is most likely
43 caused by the edge effects of upscaling a coastal gridcell's small fraction of terrestrial
44 area where high productivity occurs in a small plot, to the grid cell as a whole. By
45 evaluating NPP we are also evaluating at a secondary level litter production, which is at
46 a third level a major component of DOC production.

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48 Taken as a whole, gross primary production (GPP) was performed under simulations by
49 four PFT groups, with the largest basin-wide bulk contributions coming from boreal

50 needleleaf summer-green trees and C3 grasses (SI, Fig. S5a), the highest GPP uptake
 51 rates (3 TgC pixel⁻¹ yr⁻¹) generated by boreal needleleaf evergreen trees, and the
 52 remainder of GPP contributed by Boreal broad-leaved summer-green trees (SI, Fig. S5a).
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54 Soil respiration rates, of combined soil heterotroph and plant root respiration in our
 55 Control simulation, averaged 208 gC m⁻² yr⁻¹ (0.57 gC m⁻² d⁻¹) over the Lena basin over
 56 the period 1990-2000, which is somewhat higher than those found by Elberling (2007)
 57 in tundra soils over Svalbard, of 103-176 gC m⁻² yr⁻¹ (0.28-0.48 gC m⁻² d⁻¹). Sawamoto, et
 58 al. (2000) measured in situ summertime soil respiration over the central Lena basin and
 59 found rates of 1.6-34 gC m⁻² d⁻¹, while Sommerkorn (2008) observed rates of 0.1-3.9 gC
 60 m⁻² d⁻¹ at higher latitudes, these appearing to vary with vegetation and fire history,
 61 water table depth and temperature. Mean heterotrophic respiration rates of 1.6 gC m⁻²
 62 d⁻¹ are simulated here during July and August, in the range 0.05-2.2 gC m⁻² d⁻¹ for each of
 63 the above PFT groups. The spatial distribution of, and difference in respiration rates
 64 between PFT groups largely mirrors those for NPP (SI Fig. S5c), with maximum rates of
 65 1.4 gC m² d⁻¹ over forested sites, versus a maximum of 2.2 gC m² d⁻¹ over
 66 grassland/tundra sites (SI, Fig. S5b).
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68 Aggregated over the basin, results show that increases over the course of the 20th
 69 Century were simulated for NPP, GPP, River Discharge, DOC, CO_{2(aq.)}, autotrophic and
 70 heterotrophic respiration and CO₂ evasion, with percentage changes in the last versus
 71 first decade of +25%, +27%, 38%, +73%, +60%, +30%, +33% and +63%, respectively.
 72 (Fig. S7). It thus appears that rising temperatures and CO₂ concentrations
 73 disproportionately favoured the metabolism of carbon within the soil and its
 74 transport and mineralisation within the water column, fed by higher rates of primary
 75 production and litter formation as well as an accelerated hydrological cycle.
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80 **Table S1:** Data type, name and sources of data files used to drive the model in the study
 81 simulations.
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Data Type	Name	Source
Vegetation Map	ESA CCI Land Cover Map	(Bontemps et al., 2013)
Topographic Index	STN-30p	(Vorosmarty et al., 2000)
Stream flow direction	STN-30p	Vörösmarty et al., 2000
River surface area		(Lauerwald et al., 2015)
Soil texture class		(Reynolds et al., 1999)
Climatology	GSWP3 v0, 1 degree	http://hydro.iis.u-tokyo.ac.jp/GSWP3/
Potential floodplains	Multi-source global wetland maps	(Tootchi et al., 2019)
Poor soils	Harmonized World Soil Database map	(Nachtergaele, Freddy, Harrij van Velthuisen, Luc Verelst, N. H. Batjes, Koos Dijkshoorn, V. W. P. van Engelen, Guenther Fischer, Arwyn Jones, 2010)
Spinup Soil Carbon Stock	20ky ORCHIDEE-MICT soil carbon spinup	(Guimberteau et al., 2018)

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 85 **Table S2:** Literature sources for empirical evaluation of model output.
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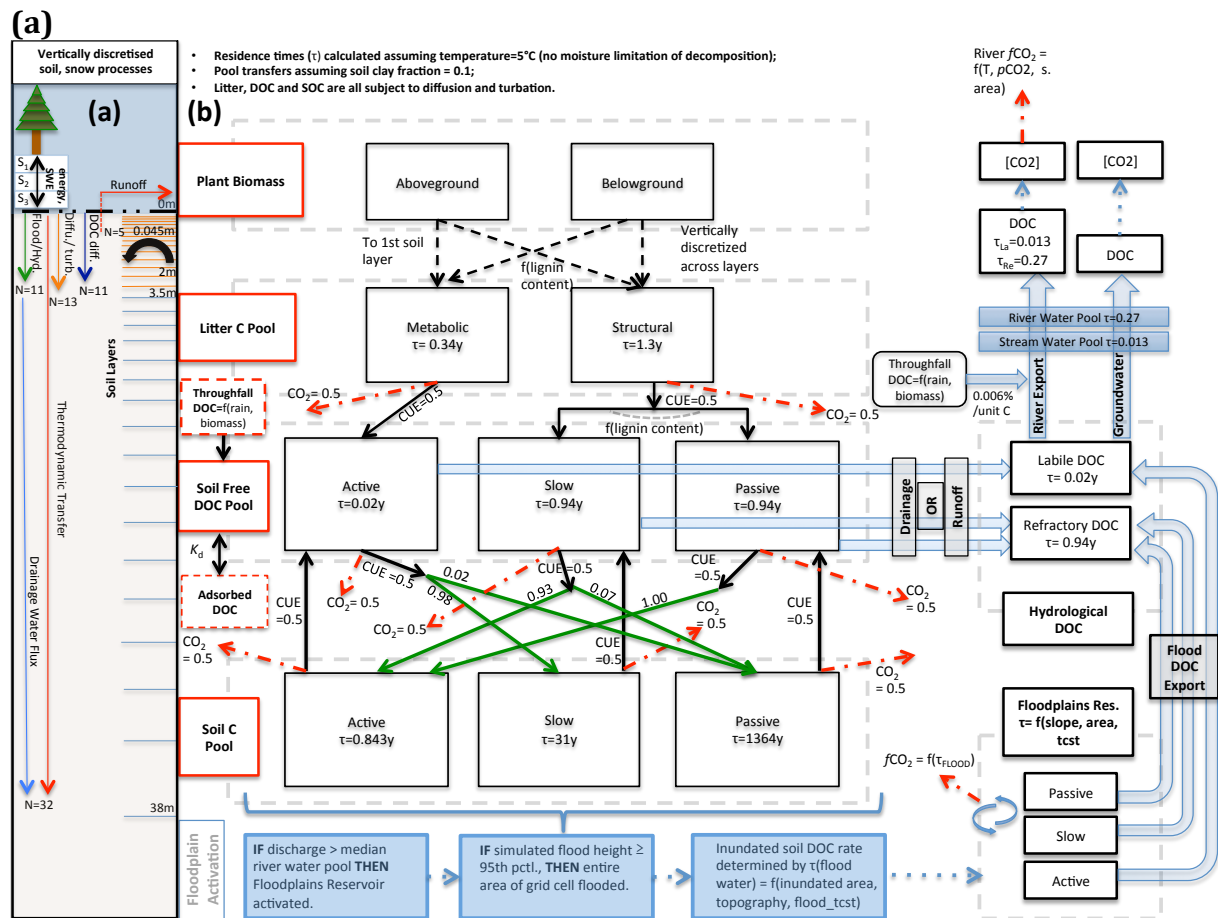
Empirical Evaluation Sources	
DOC Discharge	Cauwet and Sidorov (1996); Dolman et al. (2012); Holmes et al. (2012); Lara et al. (1998); Raymond et al. (2007); Semiletov et al. (2011); Kutscher et al. (2017).
Water Discharge	Ye et al. (2009); Lammers et al. (2001)
DOC concentration	Shvartsev (2008); Denfeld et al. (2013); Mann et al. (2015); Raymond et al. (2007); Semiletov et al. (2011); Arctic-GRO/PARTNERS (Holmes et al., 2012)
NPP	Beer et al. (2006); Lloyd et al. (2002); Roser et al. (2002); Schulze et al. (1999); Shvidenko and Nilsson, (2003)
Soil Respiration	Elberling (2007); Sawamoto et al. (2000); Sommerkorn (2008).
CO2 Evasion	Denfeld et al. (2013); Serikova et al. (2018).

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Table S3: Observed versus simulated DOC discharge (1998-2007), where we compare the output of two separate climatological datasets used as input to the model (GSWP3 and ISIMIP 2b). Also shown are the simulated versus observed DOC discharge for the six largest Arctic rivers (the "Big Six") and for the Pan-Arctic as a whole.

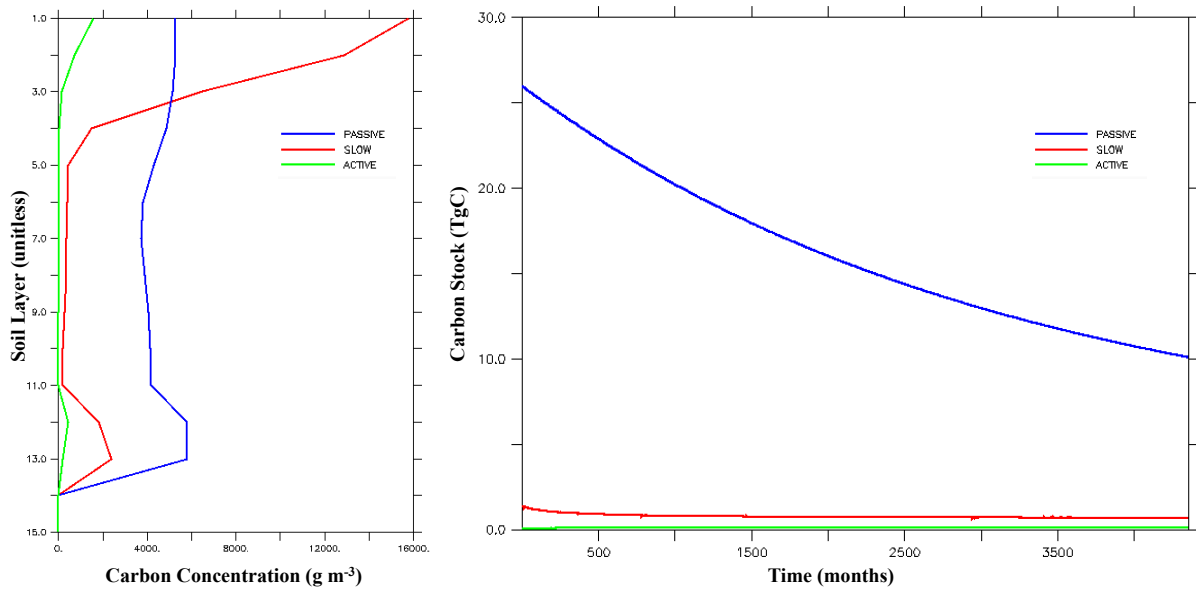
	Simulated DOC to Ocean	Simulated DOC to Ocean	Observations (Holmes et al., 2012)
	GSWP3	ISIMIP 2b	PARTNERS/Arctic-GRO
Lena	3.16	4.14	5.68
Big 6		19.36	18.11
Pan-Arctic		32.06	34.04

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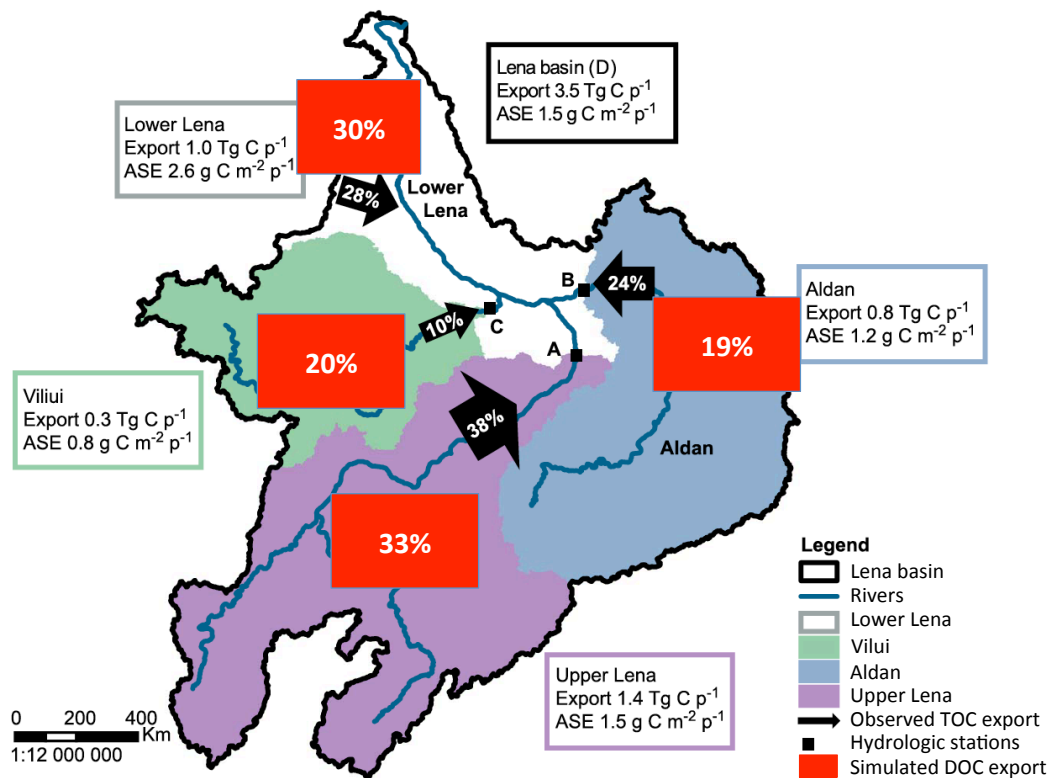


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97 (c)



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 99 **Figure S1: (a-b)** Carbon and water flux map for core DOC elements in model structure
 100 relating to DOC transport and transformation, first published in Part 1 of this study. **(a)**
 101 Summary of the differing extent of vertical discretisation of soil and snow for different
 102 processes calculated in the model. Discretisation occurs along 32 layers whose thickness
 103 increases geometrically from 0-38m. N refers to the number of layers, SWE=snow water
 104 equivalent, S_n = Snow layer n. Orange layers indicate the depth to which diffusive carbon
 105 (turbation) fluxes occur. **(b)** Conceptual map of the production, transfer and
 106 transformation of carbon in its vertical and lateral (i.e., hydrological) flux as calculated
 107 in the model. Red boxes indicate meta-reservoirs of carbon, black boxes the actual pools
 108 as they exist in the model. Black arrows indicate carbon fluxes between pools, dashed
 109 red arrows give carbon loss as CO_2 , green arrows highlight the fractional distribution of
 110 DOC to SOC (no carbon loss incurred in this transfer), a feature of this model. For a given
 111 temperature (5°C) and soil clay fraction, the fractional fluxes between pools are given
 112 for each flux, while residence times for each pool (τ) are in each box. The association of
 113 carbon dynamics with the hydrological module are shown by the blue arrows. Blue
 114 coloured boxes illustrate the statistical sequence which activates the boolean
 115 floodplains module. Note that for readability, the generation and lateral flux of
 116 dissolved CO_2 is omitted from this diagram, but is described at length in the Methods
 117 section. **(c)** (Left) Soil carbon concentrations per depth level for each soil carbon
 118 reactivity pool at the end of the spinup period. (Right) Evolution of each soil carbon pool
 119 over the course of the 400-year spinup quasi-equilibration period.
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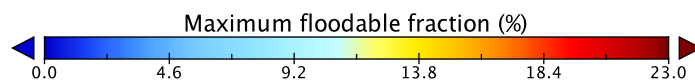
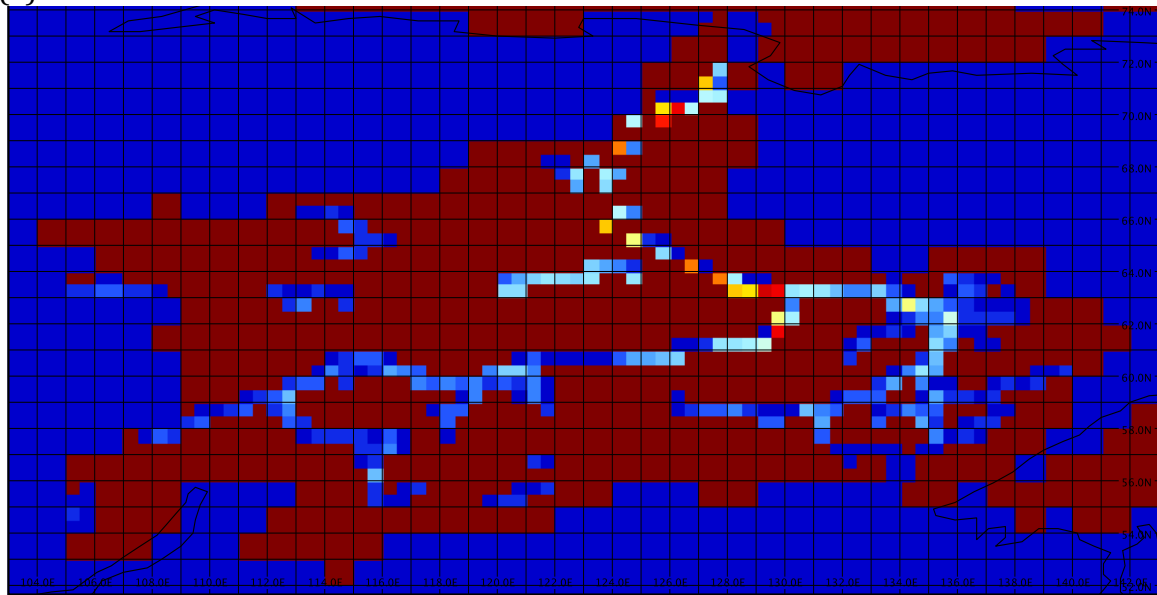


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Figure S2: Map adapted from Fig. 2 in Kutscher et al. (2017) showing proportional sub-basin contributions of TOC outflow to total TOC discharge in June and July (designated as their sampling period 'p⁻¹') of 2012-2013, as observed in Kutscher et al., 2017 (black arrows), and DOC export contributions as simulated over the period 1998-2007 by ORCHIDEE MICT-L (red boxes). Simulation pixels used in the calculation are correlates of the real-world sampling locations unless the site coordinates deviated from a mainstem hydrographic flowpath pixel –in which case a nearest 'next-best' pixel was used. Here the percentages are out of the summed mean bulk DOC flow of each tributary, not the mean DOC discharge from the river mouth, because doing so would negate the in-stream loss of DOC via degradation to CO₂ while in-stream.

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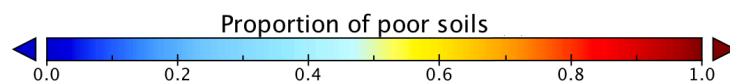
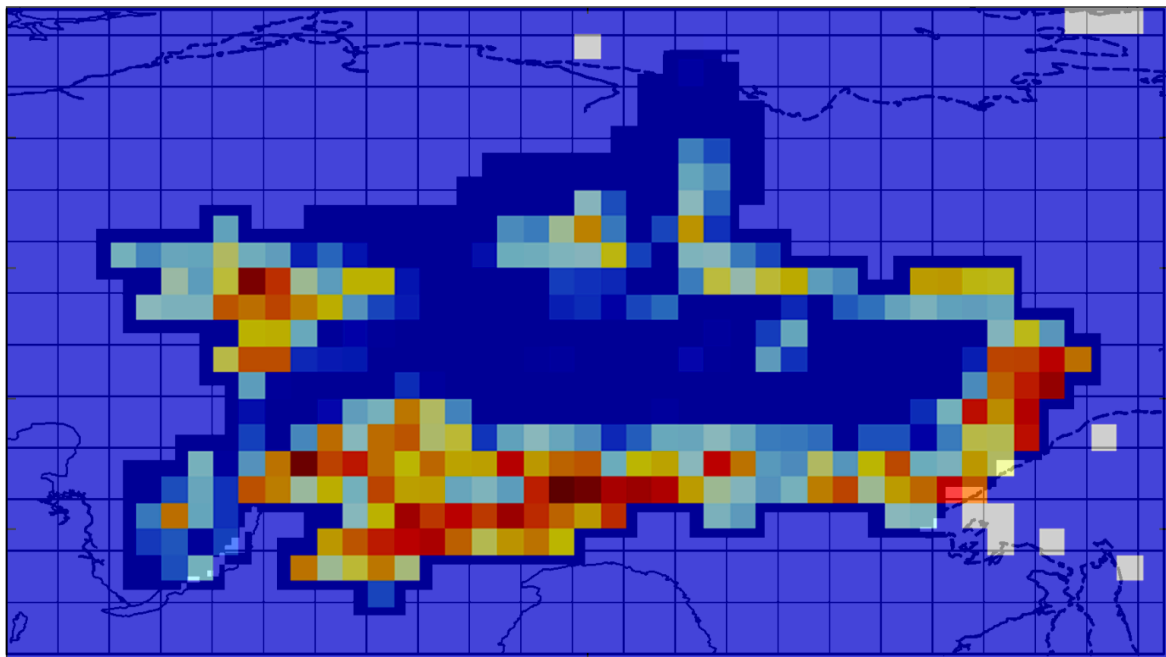
(a)



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(b)



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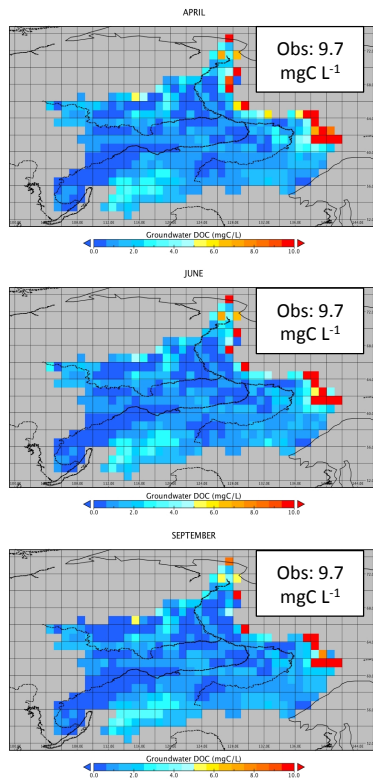
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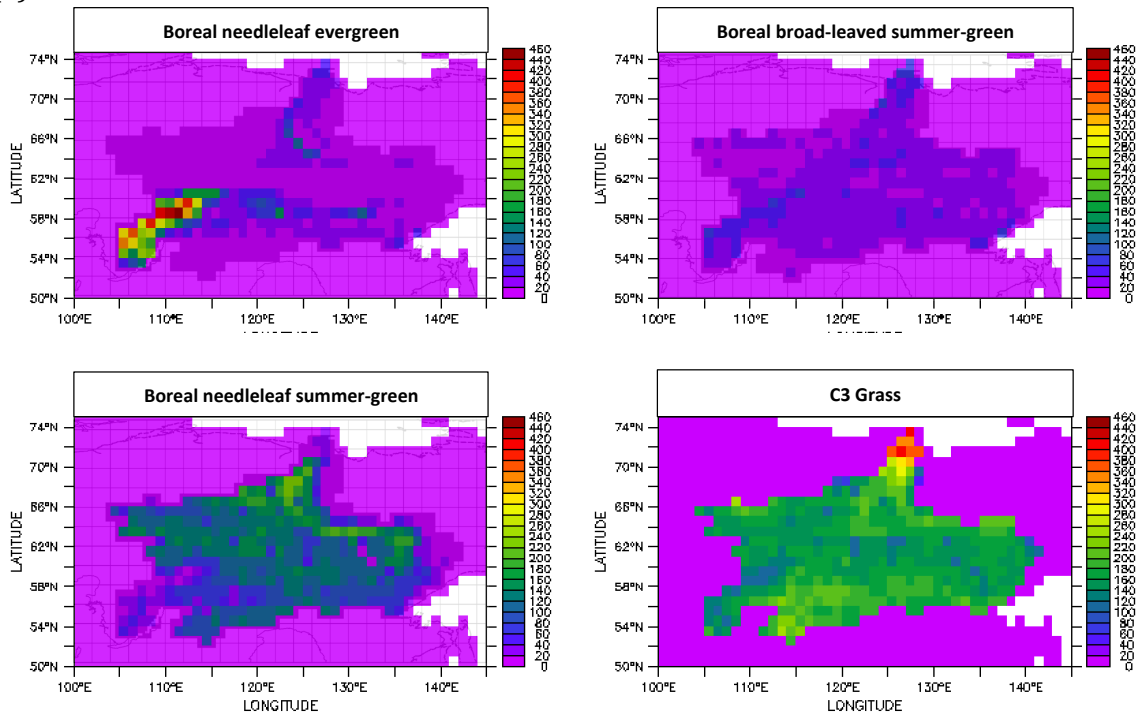
Figure S3: (a) Maximum floodable fraction of grid cells for the Lena basin per the input map from Tootchi et al. (2018). (b) Podzol and Arenosol map (Nachtergaele, Freddy, Harrij van Velthuisen, Luc Verelst, N. H. Batjes, Koos Dijkshoorn, V. W. P. van Engelen, Guenther Fischer, Arwyn Jones, 2010) used as input to the 'poor soils' module, basin mask in the background.



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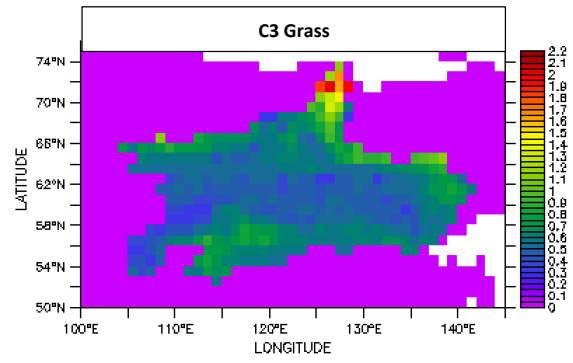
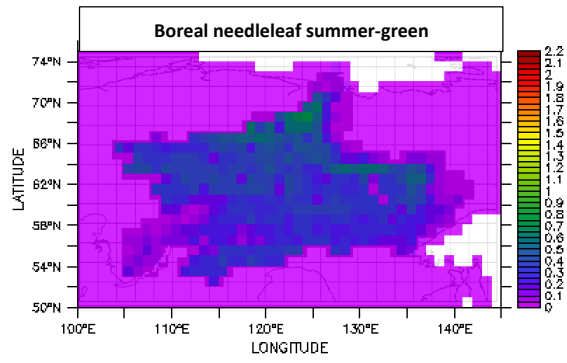
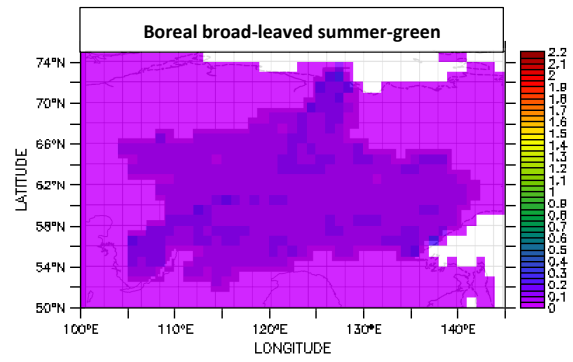
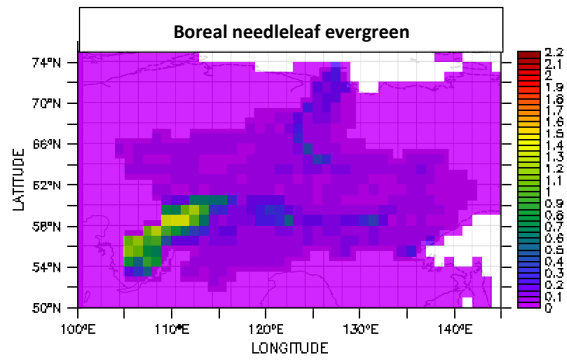
Figure S4: Groundwater DOC concentrations over the Lena basin for April, June and September averaged over 1998-2007, with mean observed concentrations for permafrost groundwater inset.

(a)



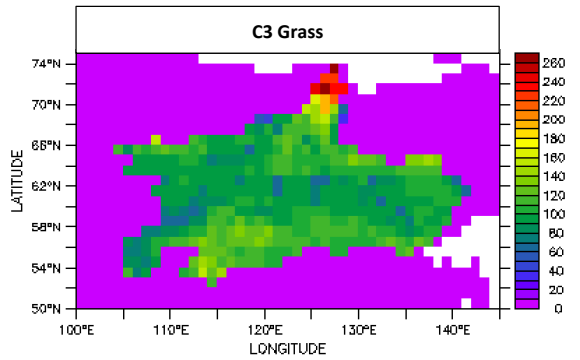
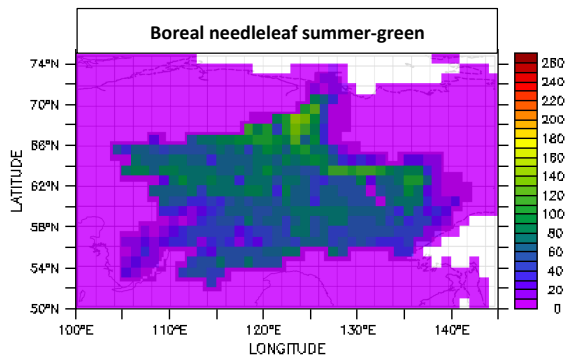
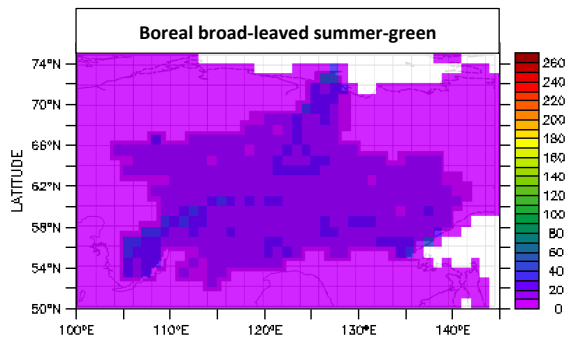
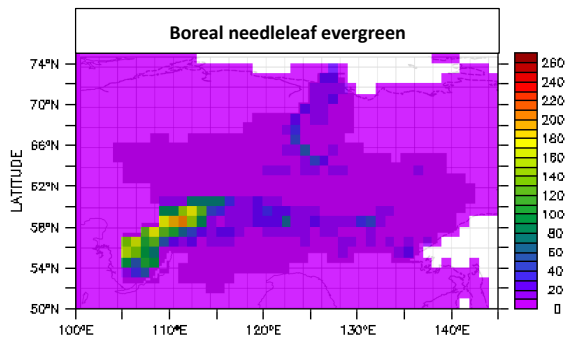
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(b)



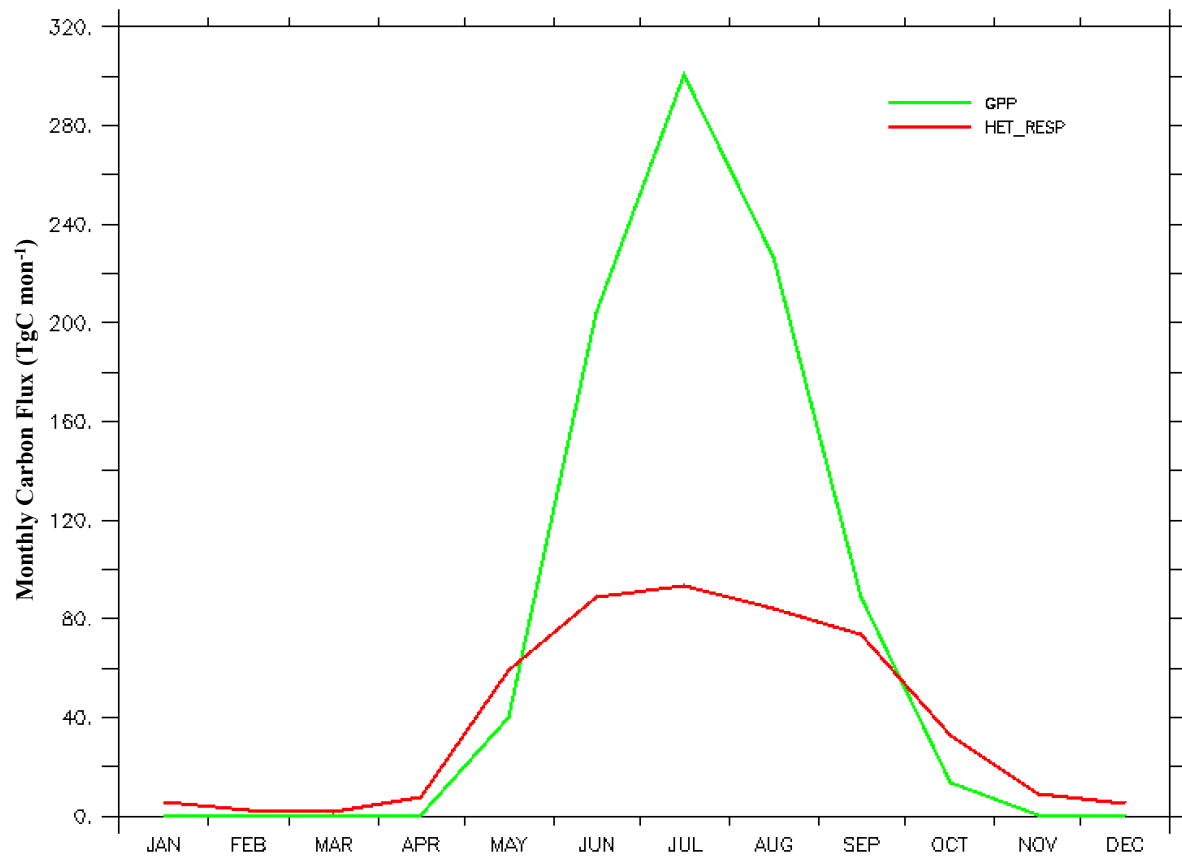
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(c)



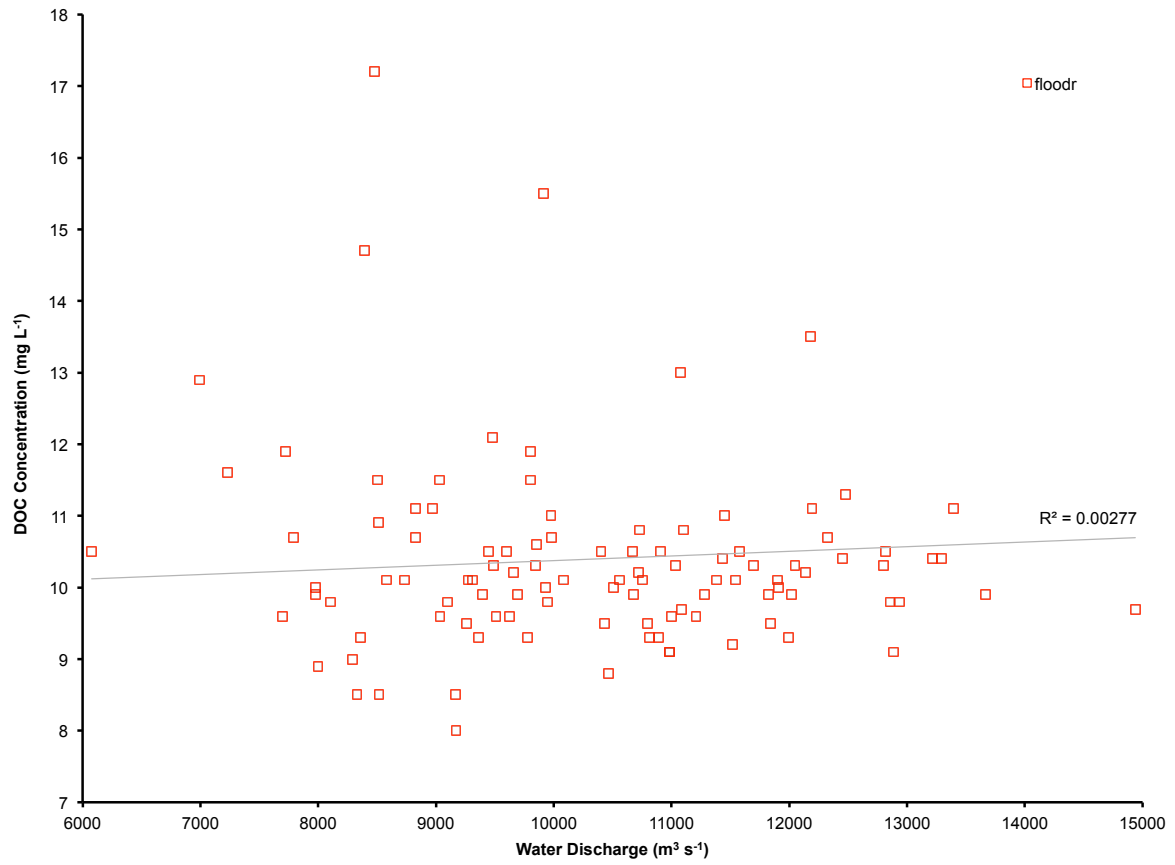
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(d)



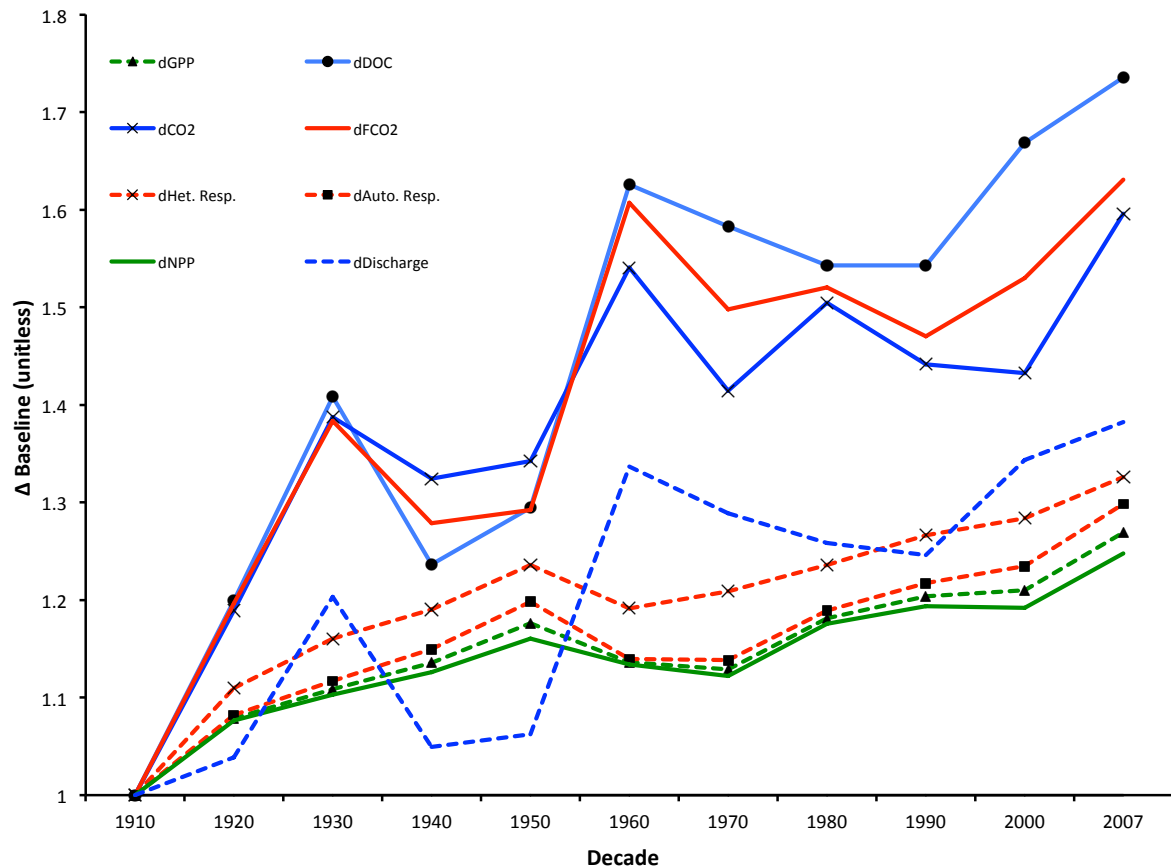
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Figure S5: (a) Absolute yearly gross primary productivity (GPP, TgC yr⁻¹) for the four relevant PFT groups over the Lena basin, averaged over 1998-2007. (b) Mean July and August soil heterotrophic respiration rates (g m² d⁻¹) for the same PFT groups as in (a), during the period 1998-2007. (c) Average yearly NPP (gC m² yr⁻¹) averaged over the period 1998-2007. All maps have the Lena basin area shaded in the background. (d) Mean monthly carbon uptake (GPP) versus its heterotrophic respiration from the soil (Het_Res) in TgC per month, over the period 1998-2007.



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Figure S6: Simulated basin-mean annual DOC concentrations (mg L⁻¹) for the floodplain water pool regressed against mean annual simulated discharge rates at Kusur (m³ s⁻¹) over 1901-2007. A linear regression with R² is plotted.



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Figure S7: Time series showing the decadal-mean fractional change in carbon fluxes normalised to a 1901-1910 average baseline (=1 on the y-axis) for NPP, GPP, autotrophic and heterotrophic respiration, DOC inputs to the water column, CO₂ inputs to the water column, CO₂ evasion from the water surface (FCO₂), and discharge.

179 **References**

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