

CORDEX-WRF v1.3: Development of a module for the Weather Research and Forecasting (WRF) model to support the CORDEX community

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Abstract. The ‘Coordinated Regional Climate Downscaling Experiment’ (CORDEX) is a scientific effort of the World Climate Research Program (WRC) for the coordination of regional climate initiatives. In order to accept an experiment, CORDEX provides experiment guidelines, specifications of regional domains and data access/archiving. CORDEX experiments are important to study climate at the regional scale, and at the same time, they also have a very prominent goal in providing regional climate data of high quality. Data requirements are intended to cover all the possible needs of stake holders, and scientists working on climate change mitigation and adaptation policies in various scientific communities. The required data and diagnostics are grouped into different levels of frequency, priority, and some of them even have to be provided as statistics (minimum, maximum, mean) over different time periods. Most commonly, scientists need to post-process the raw output of regional climate models, since the latter was not originally designed to meet the specific CORDEX data requirements. This post-processing procedure includes the computation of diagnostics, statistics, and final homogenization of the data, which is often computationally costly and time consuming. Therefore, the development of specialized software and/or code is required. The current paper presents the development of a specialized module (version 1.3) for the Weather Research and Forecasting (WRF) model, capable of outputting the required CORDEX variables. Additional diagnostic variables not required by CORDEX, but of potential interest to the regional climate modeling community, are also included in the module. ‘Generic’ definitions of variables are adopted in order to overcome model and/or physics parameterization dependence of certain diagnostics and variables, thus facilitating a robust comparison among simulations. The module is computationally optimized, and the output is divided in different priority levels following CORDEX specifications (Core, Tier1, and additional) by selecting pre-compilation flags.

This implementation of the module does not add a significant extra cost when running the model, for example the addition of the Core variables slows the model time-step by less than a 5%. The use of the module reduces the requirements of disk storage by about a 50%. The module performs neither additional statistics over different periods of time nor homogenization of the output data.

5 1 Introduction

Regional climate downscaling pursues the use of limited area models (LAM) to perform climate studies and analysis (Giorgi and Mearns, 1991). It is based on the premise that, by using LAM, modelers can simulate the climate over a region at a higher resolution as compared to the Global Climate Models (GCM). Therefore, certain aspects of the climate system can be better represented due to the higher resolution and higher complexity of parameterizations (inherent of the LAM models) used to
10 simulate physical processes which can not be explicitly resolved (e.g.: short/long-wave radiation, turbulence, dynamics of water species). This methodology has been widely used both for studying climate features, connections and processes (Jaeger and Seneviratne, 2011; Knist et al., 2014; Kotlarski et al., 2017), and to produce climate data within the scope of continental, national or regional climate change studies.

The ‘Coordinated Regional Climate Downscaling Experiment’ (CORDEX, <http://www.cordex.org/>) of the World Climate
15 Research Program (WRCP) aims to organize different initiatives devoted to regional climate all around the globe following a similar experimental design (Giorgi et al., 2009; Giorgi and Gutowski, 2015). CORDEX, with the second phase being currently under discussion, attempts to establish a series of criteria for dynamical downscaling experiments, which include setting of common domain specifications and horizontal resolutions in order to make sure that all the continental areas of the Earth are under study (e.g. in 2010 Africa was a priority and researchers worldwide volunteered to contribute with their own simulations).
20 Furthermore, CORDEX sets a series of model configurations (e.g. GCM forcing, Greenhouse Gas (GHG) evolution) to ensure that model simulations are carried out under similar conditions and therefore are inter-comparable. At the same time, CORDEX requires a list of variables necessary for a later use of model data for multi-models analysis and other climate-related research activities like climate change mitigation, adaptation and stake holders decision making policies. In order to maximize and facilitate data access, (mostly made available by the Earth System Grid Federation (ESGF), <https://esgf.llnl.gov/>), these data
25 have to be provided also following a series of homogenization criteria known as Climate and Forecast (CF) compliant (<http://cfconventions.org/>) which comes from the Coupled Model Intercomparison Project (CMIP) exercises. The list of variables required by CORDEX consists of standard model fields and some diagnostics in certain frequencies, and statistical aggregations such as minimum, maximum or mean for a given period. These variables are grouped into different priority levels (‘Core’, ‘Tier1’ and ‘Tier2’), with ‘Core’ being the mandatory list of variables (see appendix B for more details).

30 The production of these data sets is not a simple task and usually represents a big issue for the modelling community. Regional climate experiments tend to produce large amounts of data, since scientists simulate long time periods at high resolutions. Modelers have to code a software at least capable of: (1) computing a series of diagnostics, (2) concatenating model output, (3) performing statistical temporal computations and (4) producing data following CF-compliant (i.e. cmorization) criteria in

netcdf format (netCDF, Network Common Data Form <https://www.unidata.ucar.edu/software/netcdf/> a binary self-describing and machine-independent file format). Aside from being time-consuming due to its complexity and the process management, this codification also implies certain duplication of huge data-sets and additional consumption of computational resources.

Several tools (e.g. NetCdf Operators - NCO, Climate Data Operators - CDO) exist which facilitate the manipulation of netcdf files (extract/concatenate/average/join etc.), and also some other post-processing initiatives that have been made available especially to the Weather Research and Forecasting (WRF; <http://www.mmm.ucar.edu/wrf/users/>; Skamarock et al., 2008) community: WRF NetCDF Extract&Join (wrfncxnj, <http://www.meteo.unican.es/wiki/cordexwrf/SoftwareTools/WrfncXnj>), wrfout_to_cf.ncl (http://foehn.colorado.edu/wrfout_to_cf/), METtools (<https://dtcenter.org/met/users/metoverview/index.php>), and Climate Model Output Rewriter (CMOR, <https://cmor.llnl.gov/>).

WRF is a popular model for regional climate downscaling experiments. It is used world-wide in different CORDEX domains (Fu et al., 2005; Mearns et al., 2009; Nikulin et al., 2012; Domínguez et al., 2013; Vautard et al., 2013; Evans et al., 2014; Katragkou et al., 2015; Ruti et al., 2016). The model was initially designed for short-term simulations at high resolutions, but a series of modifications that had been introduced to the model code so far enhanced its capabilities, and made it appropriate also for climate experiments (Fita et al., 2010). Since WRF does not directly provide most of the required variables for CORDEX and due to the complexity of the post-processing procedures, many of the already existing WRF-climate simulations are not publicly available to the community.

This new module comes to complement the modifications introduced in the CLimate WRF (cIWRF, <http://www.meteo.unican.es/wiki/cordexwrf/SoftwareTools/CIWrf>; Fita et al., 2010). In cIWRF climate statistical values (such as minimum, maximum and mean values) of certain surface variables were introduced into the model. At the same time, evolution of Green House Gases (GHG; CO_2 , N_2O , CH_4 , $CFC-11$, $CFC-12$) can be selected from an ASCII file instead of being hard coded. Before these modifications, WRF users could only retrieve those statistical values via post-processing the standard output of the model (at a certain frequency). With the cIWRF modifications (incorporated into the WRF source code since version 3.5) statistical values are directly computed during model integration. This new CORDEX module proposes one step further by incorporating a series of new variables and diagnostics which are important for climate studies and currently WRF users can only obtain via post-processing the standard model output. At the same time, additional variables have been added into the WRF capabilities of output at pressure levels. In the current module version if the 'adaptive time-step' option is enabled in WRF, some diagnostics related to time-step selection (e.g. precipitation, sunshine duration, etc) will not be calculated properly, because there is no proper adaptation.

We present a series of modifications to the model code, and a new module (version 1.3) which will enable climate researchers using WRF to get almost all the CORDEX variables directly in the model output. With the use of this module, production of the data for regional climate purposes will become easier and faster. These modifications directly provide the required fields and variables ('Core' and almost all 'Tier1', see in appendix B for more details) during model integration, and aim to avoid the post-processing of the WRF output up to certain level. However, in this version, they do not cover all the previously mentioned aspects of the task, such as computation of statistics and the cmorization of the data.

New variables and diagnostics will be provided at the user-selected output frequency. The user still needs to post-process the data in order to obtain the different statistics required by CORDEX at daily, monthly and seasonal periods. The data cmorization can be defined as a series of processes that need to be applied to the model output in order to meet the standards provided under the CF guidelines (which follows the C-MOR standard, <https://pcmdi.github.io/cmor-site/>). These guidelines are designed to facilitate the comparison between climate models, and they represent the standard for the 'Coupled Model Intercomparison Project' (CMIP, <https://cmip.llnl.gov/>). This standardization includes the file names, variable names and metadata (units, standard name and long name), specification of geographical projections and time axis. In order to achieve a complete CF-standardization of WRF output in complete agreement with the CF requirements, substantial changes on the WRF input/output (I/O) tools would be required. This would affect backward compatibility and it has been decided to pursue in upcoming module updates. Therefore, the users of the CORDEX-WRF module will still need to perform part of the standardization. This includes joining/concatenating of WRF files, make use of standard names and attributes of the variables, file names, and finally providing the right variable with the standard attributes to describe the time coordinate.

The module also aim to establish a series of homogenization of certain diagnostics. These diagnostics can be computed following different methodologies, and consequently they may be model and/or even physical parameterization dependent. In order to avoid dependency on the model configuration (mainly sensitivity to the choice of the available different physical schemes), and to allow for a fair comparison between different simulations, a series of additional 'generic' definitions of some diagnostics are presented when possible.

The modification of the WRF model code was initiated within the development of the regional climate simulation platform from the Institute Pierre Simone Laplace (IPSL)-RegIPSL (<https://sourcesup.renater.fr/wiki/morcemed/Home>) and the CORDEX Flagship Pilot Study (CORDEX-FPS), '*Europe+Mediterranean; Convective phenomena at high resolution over Europe and the Mediterranean*', (Coppola et al., 2018) in order to obtain the variables required for the CORDEX experiment (available at: <https://www.hymex.org/cordexfps-convective/wiki/doku.php?id=protocol>) and share the code among the WRF users of the CORDEX-FPS experiment.

In this work the complete module is presented, its capabilities are demonstrated, and the results of several diagnostics are shown in order to illustrate the accuracy of the implementation. The initial section of the paper describes the modifications that have been introduced into the code followed by a description of the variables required by CORDEX. The following section demonstrates the performance tests, and gives a description of aspects which are currently missing, but will be added in the upcoming module versions. The paper finishes with a discussion and outlook section.

2 The CORDEX module

Here we present the module and we explain the modifications introduced. The steps necessary to follow in order to compile and use the module are provided as well. For a complete and detailed description of the steps to follow, the reader is referred to the wiki page of the module: <http://wiki.cima.fcen.uba.ar/mediawiki/index.php/CDXWRF> and the README file provided with

the module labeled `README.cordex`. The module has been implemented following standards of modularity which facilitates the upgrading and the introduction of new variables to it.

2.1 WRF code main characteristics

First, we provide a short description of the WRF code characteristics. The WRF model is written in Fortran 90. It is open access. It consists mainly of two parts: WPS (WRF Pre-processing System) for the preparation of the initial and boundary conditions, and the model itself. The source of the code is not fully provided. A pre-compilation process is carried out in order to automatically write certain parts of the code accordingly to a series of ASCII files and activation of certain parts of the code rely on pre-compilation flags. With the pre-compilation flags users can determine optional aspects of the model related to technical aspects of the compilation and the use of certain components like the incorporation of the ‘Community land model version 4’ (Oleson et al., 2010; Lawrence et al., 2011). Large parts of the code which are automatically written are related to the Input/Output of the model. There are a series of ASCII files provided into the `Registry` folder of the model called `registry`. These files contain the characteristics of the variables mainly: name of the variable during execution, rank and dimensions of the variable, assigned output file, name of the variable in the output file, description of the variable, and units. WRF model keeps all the variables in a Fortran pointer derived type (called `grid`). At the same time, WRF model set-up is managed through the use of a Fortran `namelist` statement which reads the ASCII file called `namelist.input` which has different sections. WRF manages the output via different streams (usually up to 23) being the standard output (`wrfout` files) the number 0. WRF model integrates the atmosphere using η as vertical variable (see more detail in, Skamarock et al., 2008) defined in equation 1 (being p_{surf} : surface pressure, p_{top} : pressure at top, p hydrostatic pressure and $\eta = 1$, surface and $\eta = 0$ on top of the atmosphere). WRF uses 3 horizontal grids (2 c-grid staggered for winds) and 2 set of vertical coordinates (1 staggered known as the ‘full’ eta levels, being the unstaggered as the ‘half’ eta levels).

$$\eta = \frac{p - p_{top}}{p_{surf} - p_{top}} \quad (1)$$

For further technical details of the model, reader is referred to WRF technical note (http://www2.mmm.ucar.edu/wrf/users/docs/arw_v3.pdf Skamarock et al., 2008) and the users’ guide web (http://www2.mmm.ucar.edu/wrf/users/docs/user_guide_v4/contents.html).

2.2 Module implementation

The module is accompanied with a new registry file called `Registry/registry.cordex` where the variables, and namelist parameters related to the module are defined. The specific set-up of the module is managed in the WRF namelist in a new section called `cordex`. Aside from the modifications of the code of the WRF model, the complete module currently consists of two new modules:

- `phys/module_diag_cordex.F`: Main module which manages the calls to the variables and performs the necessary accumulations for the calculations of statistical values (e.g.: mean, maximum, minimum)

- `phys/module_diagvar_cordex.F`: Module which contains the calculations of all the CORDEX variables separated in individual and independent 1D Fortran subroutines.

A list of detailed information on the modifications introduced is given below:

1. Main call of the CORDEX module (`module_diag_cordex.F`) has been added to
5 `phys/module_diagnostics_driver.F`, which accounts for the management of diagnostics and it has been modified in order to introduce the new pressure interpolated variables.
2. An input line to the `registry.cordex` has been added into the general `Registry/Registry.EM_COMMON`
3. The complementary pressure interpolated variables have been introduced in the related `registry Registry/registry.dia`
4. The complementary interpolated variables have been added in the module which performs the pressure interpolation
10 (`phys/module_diag_pld.F`)
5. The initialization of the modified pressure interpolation has been added in `dyn_em/start_em.F`
6. Modifications have been introduced in the `main/depend.common` and `phys/Makefile` files to get the module compiled
7. Specific changes for the inclusion of the water-budget variables have been introduced in the `dyn_em/solve_em.F`
15 module in order to get the advection terms of all water species
8. An ASCII file called `README.cordex` with the description and synthesized instructions for compilation and use is provided as well

The model output is grouped in a single file (WRF's auxiliary history output or stream #9) with a proposed file name (`auxhist9_outname` namelist parameter in the `&history` section): `wrfcordex_d<domain>_<date> regulated`
20 with the standard WRF namelist parameters of: output frequency (`auxhist9_interval`), number of time steps per file (`frames_per_auxhist9`) and format (`io_form_auxhist9`). Additional CORDEX variables required at pressure levels have been included in the WRF auxiliary output file number 23. These introduced CORDEX variables follow the file set-up via the currently existing namelist section called `diags&`.

2.3 Module use

25 Before execution of WRF some pre-processing steps are necessary to be performed by the user which encompass: compilation of the code and its specific set-up to be used during the execution time of the model and are described in the following subsections.

2.3.1 Compilation

Pre-compilation flags need to be defined by the user, depending on his/her requirements. It is necessary to keep in mind that this is done due to efficiency constraints (see below in section 6 'Optimization'), although it is not a common procedure in the standard use of WRF. Usually WRF has almost all options available from a single compilation switching options via the namelist.

Using the pre-compilation flag `CORDEXDIAG`, the CORDEX 'Core' variables will be produced. The 'Tier1' and 'additional' groups of variables can be selected via the additional pre-compilation flag `CDXWRF` (`CDXWRF=1`, for Tier1 and `CDXWRF=2` for Tier1 and the additional). The reader is referred to Appendix C for more details about the groups of CORDEX variables associated to each option. Registry file (`registry.cordex`) has to be manually modified accordingly to the selected pre-compilation flag (uncomment the associated lines).

In order to adapt this derived type to the pre-selected compilation, it is necessary also to modify the module's specific register file (`register.cordex`) according to the chosen value given to the additional pre-compilation `CDXWRF` flag (if used). This is done in a way to control the size of `grid` derived type which has a positive impact on the model performance (see below). For a complete and detailed description of these steps, reader is referred to the wiki page of the module <http://wiki.cima.fcen.uba.ar/mediawiki/index.php/CDXWRF>

According to the value given to the pre-compilation `CDXWRF` flag, different amount of variables is written out to the 'wrfcdx' output file (see more detail in appendix C):

- Using `CORDEXDIAG` and without `CDXWRF`: all the CORDEX 'Core' variables will be calculated
- `CDXWRF=1`: CORDEX 'Core' + 'Tier1' variables `clgvi`, `clhvi`, `zmla`, `[cape/cin/zlfc/plfc/lidx]{min/max/mean}`
- `CDXWRF=2`: as with `CDXWRF=1`, plus additional 3D variables at model η -level (`ua`, `va`, `ws`, `ta`, `press`, `zg`, `hur`, `hus`), 2D variables (`tfog`, `fogvisblty{min/max/mean}`, `tds{min/max/mean}`), and the water-budget variables (`wbacdiabh`, `wbacpw`, `wbacpw[c/r/s/i/g/h]`, `wbacf`, `wbacf[c/r/s/i/g/h]`, `wbacz`, `wbacz[c/r/s/i/g/h]`, `wbacdiabh{l/m/h}`, `wbacpw{l/m/h}`, `wbacpw{l/m/h}[c/r/s/i/g/h]`, `wbacf{l/m/h}`, `wbacf{l/m/h}[c/r/s/i/g/h]`, `wbacz{l/m/h}`, `wbacz{l/m/h}[c/r/s/i/g/h]`)

Moreover, the code also accounts for providing instantaneous CORDEX variables which are provided as statistics (e.g. `capemean`, `tdsmax`, or all the water-budget variables). In order to get them, user must follow certain modifications of the code (and re-compilation) in `phys/module_diag_cordex.F`, `phys/module_diagnostics_driver.F`, as well as in the registry file `registry.cordex`.

2.4 Usage

Modifications of the module include two main sets of variables: (1) new variables and diagnostics and (2) additional variables interpolated at pressure levels. These two sets of variables are provided in two separated files. New auxiliary output file in the ninth stream provides all the new variables and diagnostics required by CORDEX. Additional pressure interpolated variables

are included in the 23rd stream. Each of these files have to be set-up in the namelist in the same way as it is done with the standard WRF output files.

A new section labeled `cordex` has to be added into the WRF's namelist which allows to choose/set-up different options of the module. The description of all the available options is provided in Table 1. In this section the user is required to choose the implementation of the diagnostics to use, to provide values to some parameters for certain diagnostics, and to activate/deactivate some of the most computationally costly diagnostics. Default values for all the options are provided in order to facilitate the use of the module.

This module has been tested under different High-Performance Computing (HPC) environments and compilations. It has been compiled with two compilers: `gfortran` and `ifort`. Different parallelization paradigms: serial, distributed memory and hybrid (distributed and shared) and the parallelized version of the `netcdf` libraries. The tests have been performed mainly using 2-nested domains with the second one being at the convection permitting resolution (no cumulus scheme activated). Under all these circumstances the module worked as expected.

Since the current version v1.3 of the module, a text message with the version of the module is printed in the standard output at the first time-step of the model run in order to facilitate the detection of the module version that is being used.

15 3 CORDEX variables

CORDEX requires a series of mandatory variables grouped in the 'Core' level, and additional variables grouped in 'Tier1' and 'Tier2' levels. Furthermore, CORDEX also requires statistical values of specific variables, besides the instantaneous values. To meet the CORDEX specifications, regional climate models have to provide three kind of variables:

- instantaneous: values obtained at each model integration time-step. An instantaneous value represents the field at the given instant of time all over the given space encompassed within the grid point.
- statistics: values obtained as statistics of consecutive instantaneous values along a given period of time. The statistical computation could be: minimum, maximum, mean or accumulated value, as well as the flux. Thus, one statistical variable represents the temporal statistics of the field for a given period of time all over the given space encompassed within the grid point. CORDEX guidelines also require different temporal aggregations: 3-hourly, daily, monthly and seasonal
- fixed: values which do not have an evolution in time. These fields are fixed all over the simulation.

The WRF I/O file managing system provides an infrastructure for more than 20 different output files (called streams) at the same time. Each file is independently managed, and therefore in the namelist a user has to set-up mainly 2 different options for each output stream: frequency of an output (frequency of writing out the variables to an output file in minutes, e.g. 30, 60), number of frames per file (e.g. for 3-hourly frequency 8 frames per file will give a daily output). Variables can be written in multiple streams (selected via the `'registry'` files). During the model integration, at the given time-step corresponding to the defined output frequency, data will be written out to the output file. When a given file reaches the selected amount of

Table 1. Set-up parameters of the `module_diag_cordex` module for the WRF namelist contained in the `cordex` section. See ‘variables’ sections for more details of the meaning of each methodology. The methodologies preferred by CORDEX are marked by ^a, the ones without preference by CORDEX are marked by ^b in these cases, users can select the method according to their experience.

name & value	description	default value
<code>output_cordex = 0</code>	CORDEX Diagnostic de-activation	0
<code>output_cordex = 1</code>	CORDEX Diagnostic activation	0
<code>psl_diag = 1</code>	sea-level pressure diagnostic following hydrostatic Shuell correction (Stackpole and Cooley, 1970)	3
<code>psl_diag = 2</code>	psl diagnostic following a target pressure (Benjamin and Miller, 1990)	3
<code>psl_diag = 3</code>	psl diagnostic following ECMWF method (Yesad, 2015)	3 ^a
<code>psmooth = 5</code>	number of passes of neighbor filtering (mean of the grid point with its 8 neighbors) of psfc (only for psl_diag=2)	5
<code>ptarget = 70000.</code>	pressure [Pa] target to be used by psl_diag=2	70000.
<code>wsgs_diag = 1</code>	wind-gust diagnostic following (Brasseur, 2001)	1 ^b
<code>wsgs_diag = 2</code>	wind-gust following heavy precipitation method	1
<code>output_wb = 0</code>	deactivation of the computation of water-budget variables (Fita and Flaounas, 2018)	0
<code>output_wb = 1</code>	activation of the computation of water-budget variables	0
<code>wsz100_diag = 1</code>	wind extrapolation at z100m_wind using power-law method	1 ^b
<code>wsz100_diag = 2</code>	wind extrapolation at z100m_wind using logarithmic-law method	1
<code>wsz100_diag = 3</code>	wind extrapolation at z100m_wind using Monin-Obukhov method	1
<code>z100m_wind = 100.</code>	height [m] to extrapolate winds for wsz100_diag	100
<code>zmlagen_dqv = 0.1</code>	percentage of variation of mixing ratio to determine mixed layer depth used in zmlagen computation (Nielsen-Gammon et al., 2008)	0.1 ^b
<code>zmlagen_dtheta = 1.5</code>	increment in K of potential temperature from its minimum within the MLD used in zmlagen computation	1.5 ^b
<code>potevap_diag = 1</code>	potential evapotranspiration using bulk computation (Manabe, 1969)	2 ^b
<code>potevap_diag = 2</code>	potential evapotranspiration using Milly92 correction (Milly, 1992)	2
<code>convxtrm_diag = 0</code>	deactivation of diagnostic of extremes from convection indices	0
<code>convxtrm_diag = 1</code>	activation of diagnostic of extremes from convection indices	0
<code>fogvisibility_diag = 1</code>	diagnostic of visibility inside fog (Kunkel, 1984)	3 ^b
<code>fogvisibility_diag = 2</code>	RUC method (Smirnova et al., 2000)	3
<code>fogvisibility_diag = 3</code>	FRAML 50% prob (Gultepe and Milbrandt, 2010)	3
<code>fogvars = 1</code>	use 3D variables (hur, closest level to surface) to diagnose fog	1 ^b
<code>fogvars = 2</code>	use sfc variables (hurs) to diagnose fog (not available for fogvisibility_diag = 1)	1

^apreferred by CORDEX

^bNo preference is specified by CORDEX

frames, it is closed and a new one is open. The file name usually follows a criteria of a given header name (e.g. 'wrfcdx' for this module) and the current date of the simulation which is also set-up in the namelist.

The CORDEX-WRF module is designed to provide the variables using the 9th stream, without reducing any of the capabilities of the model. Following this criteria, the module uses the same structure and components of the model designed to manage its I/O. This means that the statistical values are directly provided using the internal values between output frequencies. This ensures that, for example, a minimum value would exactly be the minimum value that the model simulated between output times. These variables are re-initialized after each stream output time (see figure 1). The WRF model is used in a myriad of applications and regions, thus it was decided that statistics values will be provided at the selected frequency of the 9th stream. This gives more flexibility, allowing a user to get e.g. high frequency outputs. However, this will require the user to perform a post-processed aggregation of the output files in order to provide the required CORDEX statistics at the 3-hourly, daily, monthly and/or seasonal periods. Users are strongly encouraged to use the output frequencies for the 9th stream which are easy to combine in order to retrieve the required CORDEX statistics. It is necessary to highlight that the statistics for a given period contained in the 9th stream corresponds just to the instant time of writing the field into the file (e.g. on a 3-hour output frequency the value inside the file at [HH]+3:00:00 represents the statistics from [HH]:00:00 to [HH]+2:59:59). The WRF I/O does not allow to produce static/fixed fields, therefore this group of variables are not provided by the module.

3.1 Generic methodology

The list of variables requested by the CORDEX experiment (see the Table C in appendix) are intended to be useful for the climate change mitigation, adaptation and decision making communities. Note that CORDEX-FPS might require other variables or requires some of them at different frequency of output in comparison to a standard CORDEX requested list of variables. Taking into account the performance of the model, variables are computed at specific frequencies: (1) at all time steps when a statistic value (accumulation/flux, minimum, maximum and/or mean) of the variable is required, (2) at the given time step, when a variable that is used for the diagnostic is updated following the configuration from the namelist (e.g. cloud derived variables, and the frequency of activation of the radiation scheme), (3) instantaneous values which are computed only at the time step when the output is written out (see the figure 1 for more details).

The list of the added variables to the existing direct level pressure interpolated output is provided in Table 2 and are gathered in the 23rd auxiliary WRF output file with the standard name `wrfpres_d<domain>_<date>` (in WRF's namelist notation). At the same time, in order to avoid overloading the execution of WRF, the section of the code with the pressure interpolation has been also modified. Now the interpolation is only computed at the time-step coincident with the output frequency (selected also in the namelist as well as the characteristics of the pressure interpolation).

The different statistical values are initialized (as shown in figure 1) at the first time step after output time. More details on how certain diagnostic variables have been integrated, and implemented in WRF are provided in the following sections. Furthermore, a series of plots accompanying different definitions of the diagnostic variables are presented as well. The intention of these figures is to illustrate the consistency of the implemented diagnostics. These preliminary outcomes are not for validation purposes, but rather to show that the diagnostic variable have been correctly introduced. A complete analysis in order to find the

Figure 1. Calculation of diagnostics according to the time step for each kind of variable (see Tables in appendix C)

compute variables required at each time step [*a* and *s* kind variables]

IF (*it* == *schemefreq*)

 compute variable related to the scheme [*t* kind variables]

ELSE

 keep value of the previous time step [*t* kind variables]

END IF

IF (MOD(*it*,*outputfreq*) == 0)

 initialize statistics values [*s* kind variables]

$\chi_{flux} = \chi(it)$

$\chi_{min} = \chi(it)$

$\chi_{max} = \chi(it)$

$\chi_{mean} = \chi(it)$

ELSE

 compute statistics values [*s* kind variables]

$\chi_{flux} = \chi_{flux} + \chi(it)$

 IF ($\chi(it) < \chi_{min}$) → $\chi_{min} = \chi(it)$

 IF ($\chi(it) > \chi_{max}$) → $\chi_{max} = \chi(it)$

$\chi_{mean} = \chi_{mean} + \chi(it)$

END IF

IF (*output_time*)

 compute variables required only at output time [*o* kind variables]

 finish certain statistics [*s* kind variables]

$\chi_{flux} = \chi_{flux} / (N_{steps} \times dt)$

$\chi_{mean} = \chi_{mean} / N_{steps}$

END IF

Table 2. Description of CORDEX additional pressure interpolated variables provided with the module

CF name	WRF name	description	units
hus	HUS_PL	specific humidity	1
wa	W_PL	vertical wind speed	m.s-1
ua	UER_PL	Earth-rotated wind x-component	ms-1
va	VER_PL	Earth-rotated wind y-component	ms-1
ws	WS_PL	wind speed	ms-1

most accurate methodology for the calculation of a certain diagnostic would require devoted climate simulations, and enough observations to validate them. Such validation is out of the scope of this study. We do select ‘more appropriate’ options based on the experience within the scientific community or according to the CORDEX specifications. These options are set as the ‘default’ options within the namelist.

- 5 One should be aware that certain diagnostics use variables for their calculation which might only be available when specific physical schemes are selected. When this happens, zero values are returned. This undesired outcome is, when possible, fixed by using a ‘generic’ definition of the diagnostics.

3.2 Core variables

These are the basic variables required by CORDEX. Most of them are standard fields and therefore tend to require simple calculations from the currently available variables from the WRF model. These variables are obtained by setting the pre-compilation flag CORDEXDIAG, and will appear in two different files: 3D variables at pressure levels (the WRF model internally interpolate them since it uses the η coordinate in the vertical) will appear in the output file with the 23rd stream (mainly called `wrfpress`), and the 2D variables in the module’s output file `wrfcdx`.

3.2.1 3D at pressure-levels

- 15 These are the additional variables which have been added into the WRF pressure-level integration module. Their values will be written in the 23rd output stream in addition to the ones currently available. All of them are instantaneous values.

hus: humidity

3D atmospheric specific humidity (`hus`)¹ and relative humidity (`hur`) are computed at the unstaggered model ($half-\eta$) levels. Specific humidity is simply obtained from water vapor mixing ratio using equation 2 (named *QVAPOR* in WRF). Relative humidity can be obtained following the Clausius-Clapeyron formula and its approximation from the well-known *August-Roche-Magnus* formula for saturated water vapor pressure e_s (equations 3, 4 and 5)

¹from the AMS glossary http://glossary.ametsoc.org/wiki/Specific_humidity

$$hus = \frac{QVAPOR}{QVAPOR + 1} \quad (2)$$

$$e_s = 6.1094 * e^{\frac{17.625 * tempC}{tempC + 243.04}} \quad (3)$$

$$w_s = \frac{0.622 * e_s}{presshPa - e_s} \quad (4)$$

$$hur = \frac{QVAPOR}{w_s * 1000}. \quad (5)$$

- 5 with $tempC$: being temperature in degree Celsius ($^{\circ}C$), $presshPa$: pressure (hPa), e_s : saturated water vapor pressure (hPa), w_s : saturated mixing ratio ($kgkg^{-1}$).

press: air-pressure

WRF model integrates the perturbation of the pressure field from a reference one. Thus to obtain the full pressure at unstaggered model η levels, the user is required to combine two different fields as shown in equation 6,

$$10 \quad press = P + PB \quad (6)$$

where PB : WRF base pressure (Pa), P : WRF perturbation pressure (Pa).

ta: air-temperature

This variable states for the 3D atmospheric temperature on unstaggered model η levels. WRF model equations are based on the perturbation of potential temperature, therefore a conversion to actual temperature is required, and it is performed as indicated

- 15 by equation 7,

$$ta = (T + 300) \left(\frac{P + PB}{p0} \right)^{R/C_p} \quad (7)$$

where T : WRF 3D temperature output (which is as potential temperature perturbation from the base value, which in WRF equals to 300 K), $p0$: the pressure reference (100000 Pa).

ua/va: Earth rotated wind components

- 20 These variables state for the 3D atmospheric wind components following Earth coordinates on unstaggered model η levels. WRF model equations use the Arakawa-C horizontally staggered grid with wind components following the grid direction. In order to get actual winds following the Earth geographical coordinates, a transformation shown in equation 8 is required,

$$\begin{cases} U_{unstg}(1 : dimx, 1 : dimy) = 0.5[U_{stg}(1 : dimx - 1, 1 : dimy) + U_{stg}(2 : dimx, 1 : dimy)] \\ V_{unstg}(1 : dimx, 1 : dimy) = 0.5[V_{stg}(1 : dimx, 1 : dimy - 1) + V_{stg}(1 : dimx, 2 : dimy)] \\ \begin{cases} ua = U_{unstg}cosa - V_{unstg}sina \\ va = U_{unstg}sina + V_{unstg}cosa \end{cases} \end{cases} \quad (8)$$

where U_{unstg} : unstaggered WRF eastward wind (ms^{-1} , $[1, dimx]$), V_{unstg} : unstaggered WRF northward wind (ms^{-1} , $[1, dimy]$), U_{stg} : x-staggered WRF eastward wind (ms^{-1} , $[1, dimx+1]$), V_{stg} : y-staggered WRF northward wind (ms^{-1} , $[1, dimy+1]$), $cosa$: local cosine of map rotation (1), $sina$, local sine of map rotation (1).

zg: geopotential height

- 5 As in the case of air-pressure, WRF model also integrates the perturbation of the geopotential field from a reference or base one. Thus to obtain the full geopotential height on staggered model η levels, the user is required to combine the two WRF fields and it is also de-staggered as it is shown in equation 9,

$$zg_{staggered} = PH + PHB \quad (9)$$

$$zg(k) = 0.5(zg_{staggered}(k) + zg_{staggered}(k+1)), k = [1, dz]$$

- 10 where PHB : WRF base geopotential height ($m^2 s^{-2}$), PH : WRF perturbation of the geopotential height ($m^2 s^{-2}$), $zg_{staggered}$: staggered geopotential height $k = [1, dimz + 1]$, zg : un-staggered geopotential height $k = [1, dimz]$

3.2.2 2-dimensional

- Here is provided the list of the the added 2-dimensional CORDEX variables. Some of them are diagnosed as a combination of 3-dimensional variables, some are required as instantaneous values, and others as statistics. The fact that the module provides 15 2D variables online using 3D fields, shows one more key advantage of the module related to disk space. Thanks to these online calculations when using the module, a user does not need anymore to store large amount of 3D data from the model in order to post-process them. This reduces the requirements of disk space by a factor of around 2.

pr, prc, prl, prsh, prsn: precipitation fluxes

- The total precipitation flux (pr) is computed as the sum of all types of precipitation fields in the model accumulated along the 20 9th stream output frequency (9freq) divided by this period of time (9freq), as it is shown in equation 10,

$$pr = \frac{\sum_{it}^{9freq} RAINCV(it) + RAINNCV(it) + RAINSHV(it)}{Nsteps \times \delta t} \quad (10)$$

where $RAINCV$: instantaneous precipitation from cumulus scheme (kgm^{-2}), $RAINNCV$: instantaneous precipitation from microphysics scheme (kgm^{-2}), $RAINSHV$: instantaneous precipitation from shallow-cumulus scheme (kgm^{-2}), $Nsteps$: number of time steps, and δt : time step length (s) to achieve the 9th stream frequency output time ($9freq = Nsteps \times \delta t$).

- 25 In this version, the computation of the accumulated values does not take into account configurations of the model with adaptive time-step. When adaptive time-step is used, we strongly discourage the use of these variables.

Each individual precipitation flux is also provided as:

$$\left\{ \begin{array}{l} prc = \frac{\sum_{it}^{9freq} RAINCV(it)}{Nsteps \times \delta t} \\ prl = \frac{\sum_{it}^{9freq} RAINNCV(it)}{Nsteps \times \delta t} \\ prsh = \frac{\sum_{it}^{9freq} RAINSHV(it)}{Nsteps \times \delta t} \end{array} \right. \quad (11)$$

Solid precipitation flux (*prsn*) only accounts for frozen precipitation. Depending on the selected micro-physics scheme chosen in the `'namelist'`, this variable might account for the precipitation of snow, graupel and hail. It is computed as it is shown in equation 12,

$$prsn = \frac{\sum_{it}^{9freq} prins(it) \times SR(it)}{Nsteps \times \delta t} \quad (12)$$

where *prins*: instantaneous total precipitation (kgm^{-2} , previously obtained), *SR*: fraction of solid precipitation (% , variable provided by WRF).

Radiative flux

Surface upwelling shortwave radiation flux (*rsus*, Wgm^{-2}) and surface upwelling longwave radiation flux (*rlus*, Wgm^{-2}) are understood as the shortwave and longwave radiation from Earth's surface. They are directly provided by radiation schemes CAM² and RRTMG³ (`sw_ra_scheme = 3, 4`) as instantaneous variables `swupb` and `slupb`. When there is no use of such schemes, it is recommended to use the 'generic' definition instead (`rsusgen`, `rlusgen`, see in next section). Statistical retrieval for the surface fluxes follows the same methodology as for the precipitation fluxes.

Outgoing radiative fluxes at top the atmosphere are also provided being `'rsut'` for mean Top of the Atmosphere (TOA) outgoing shortwave radiation (in Wgm^{-2}) and `'rlut'` for longwave. However there is not a 'generic' implementation of these variables.

sund: duration of sunshine

This variable accounts for the the sum of the time for which the direct solar irradiance (downwelling short-wave radiation, *rsds*) exceeds $120 Wm^{-2}$ (WMO, 2010a) implemented following equation 13 and provided in seconds. In order to provide an example of the correct implementation of this diagnostics preliminary results are shown in figure 2. The figure shows the 'sund' values and compare them with the incoming solar radiation. It is shown how the 'sund' values vary accordingly to the moment of the day with zero values during night (left panel) or persistent totally cloud covered regions (map at the right panel)

²Community Atmosphere Model

³Rapid Radiative Transfer Model

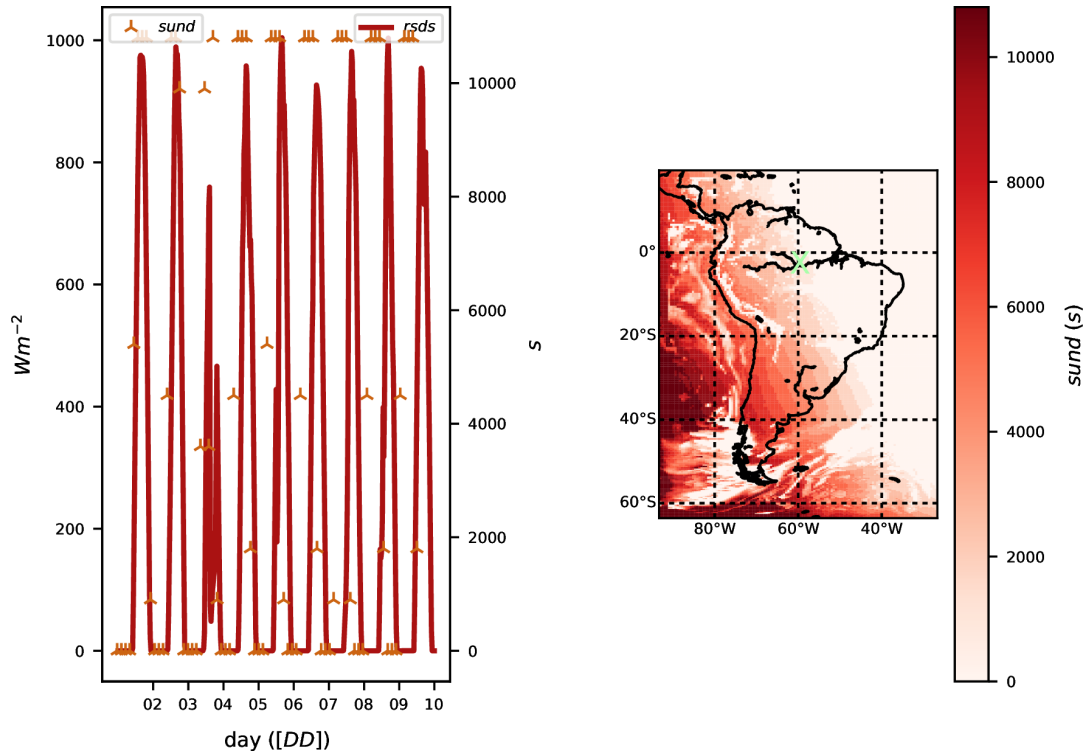


Figure 2. Temporal evolution at $S 62^{\circ} 4' 38.00''$, $4^{\circ} 58' 55.51'' W$ (left panel) of shortwave downward radiation ($rsds$, red line, left y-axis) and sunshine duration for a 3-hourly 9freq ($sund$, stars, right y-axis). Sunshine length map on 2012-12-09 at 15 UTC. X denotes position of the temporal evolution

$$sund = \sum_{it}^{9freq} \delta t [SWDOWN(it) \geq 120 Wm^{-2}] \quad (13)$$

where $SWDOWN$: downward shortwave radiation (Wm^{-2}), δt : time-step length (s).

tauvv: surface downward wind stress

Instantaneous surface downward wind stress at 10m accounts for the force that winds exerts on the Earth's surface. It is implemented following the equation 14

$$tauv = (C_D uas^2, C_D vas^2) \quad (14)$$

where, C_D : drag coefficient (1), uas : Earth-rotated eastward 10 m surface wind (ms^{-1}), vas : Earth-rotated northward 10 m surface wind (ms^{-1}). The drag coefficient is non-zero only for certain options of the surface layer physics (`sf_sfclay_physics`)

parameter in the namelist): 1 (MM5-similarity), or 5 (MYNN surface layer). A ‘generic’ formulation has been introduced when these schemes are not used.

psl: sea level pressure

This variable accounts for the instantaneous pressure extrapolated to the sea level. It represents the value of the pressure without the presence of orography. In order to provide a framework ready to implement different methodologies, three different methods have already been implemented. The choice of the method can be controlled by a new `namelist.input` parameter labeled `psl_diag` in `cordex` section. The implemented methods are:

- The hydrostatic-Shuelli method (Stackpole and Cooley, 1970) already implemented in the the module `phys/module_diag_afwa.F`, assuming a constant lapse-rate of -6.5 K km^{-1} , selected when in WRF ‘cordex’ namelist section setting the parameter `[psl_diag = 1]`
- The ‘ptarget’ method (Benjamin and Miller, 1990) that uses smoothed surface pressure and a target upper-level pressure, already implemented in the WRF post-processing tool called `p_interp.F90` `[psl_diag = 2]`
- The ECMWF method (Yesad, 2015) taken from the Laboratoire de Météorologie Dynamique GCM (LMDZ; Hourdin et al., 2006) from the module `pppmer.F90`, following the methodology by Mats Hamrud and Philippe Courtier from ECMWF `[psl_diag = 3]`

According to the CORDEX specifications, the default method is the ECMWF method. When choosing the ‘ptarget’ method (`psl_diag = 2`), also the degree of smoothing of the surface using the surrounding nine point average can be chosen by selecting a number of smoothing passes (`psmooth`, default 5), and the upper pressure that has to be used as the target (`ptarget`, default 700 *hPa*).

Figure 3 shows the different outcomes applying each method. There are some problems with the ‘ptarget’ method in both `psl` estimates (mountain ranges can still be inferred) and borders for each parallel process (lines in figures showing differences among methods) when the spatial smoothing is applied. Lines showing the limits of the parallel processes appear because one can not obtain the proper values from outside the correspondent tile of the domain associated with each individual parallel process.

Cloud derived variables

Four cloud derived variables are required by CORDEX: the total cloudiness (`clt`) and the cloudiness for each grid point at three different vertical layers above ground (low: $p \geq 680 \text{ hPa}$, labeled `cll`; medium: $680 < p \leq 400 \text{ hPa}$, `clm`; high: $p < 400 \text{ hPa}$, `clh`). These cloud diagnostics are provided as mean values.

The module computes these variables taking the cloud fraction of a given grid cell and level as input. The cloud fraction in WRF is computed by the radiative scheme, and it is called at a frequency given by `radt` parameter in WRF namelist. Due to the large computational cost of the radiative scheme, `radt` usually is larger than the time-step of the simulation. This determines when cloud fraction is also actualized to meet the evolved atmospheric conditions. Cloud fractions can be computed in the model using different methodologies. It would be possible to make available these methodologies as another choice in

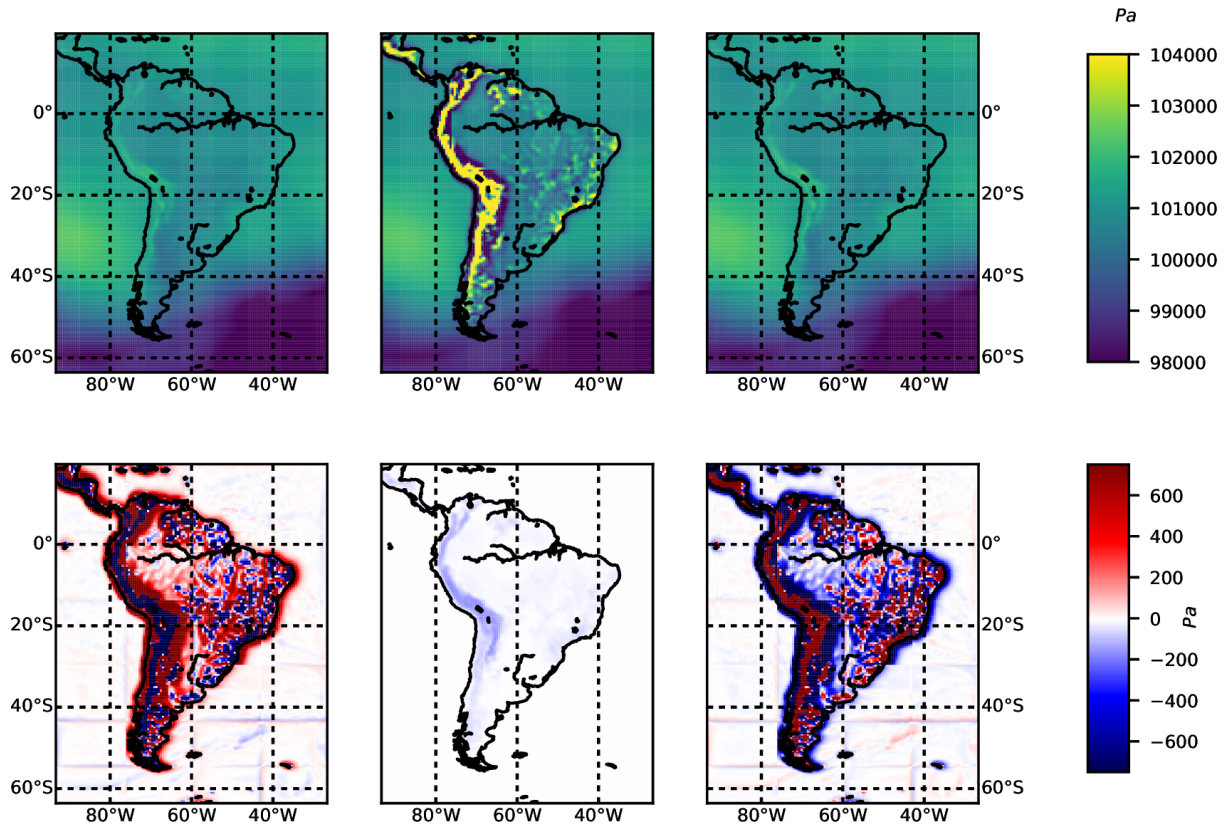


Figure 3. Sea level pressure estimates following the hydrostatic-Shueller method at a given time step ($psl^{shueller}$, upper left), p_{target} ($psl^{ptarget}$, upper middle) and ECMWF (psl^{ecmwf} , upper right). Bottom panels show differences between methods $psl^{shueller} - psl^{ptarget}$ (bottom left), $psl^{shueller} - psl^{ecmwf}$ (bottom middle) and $psl^{ptarget} - psl^{ecmwf}$ (bottom right)

the namelist section and then compute the cloud fraction at each time-step. However, in order to be consistent with the radiative cloud effects that the simulation is experiencing, this method was discarded. Thus, the cloud values provided by the module follow the same frequency of refreshing rate as the one set for radiation in the namelist level (τ_{adt} value).

- The most common implementation of ‘clt’ found in other models (in particular most GCMs) assumes ‘*random overlapping*’.
- 5 Random overlapping assumes that adjacent cloud layers are from the same system, hence are maximumly overlapped (Geleyn and Hollingsworth, 1979). In the module, the methodology from the GCM LMDZ was implemented. In this GCM, calculation of the total cloudiness and different layers’ cloudiness is done inside the subroutine `newmicro.f90`. The method basically consists in a product of the consecutive non-zero values of cloud fraction as it is shown in equation 15,

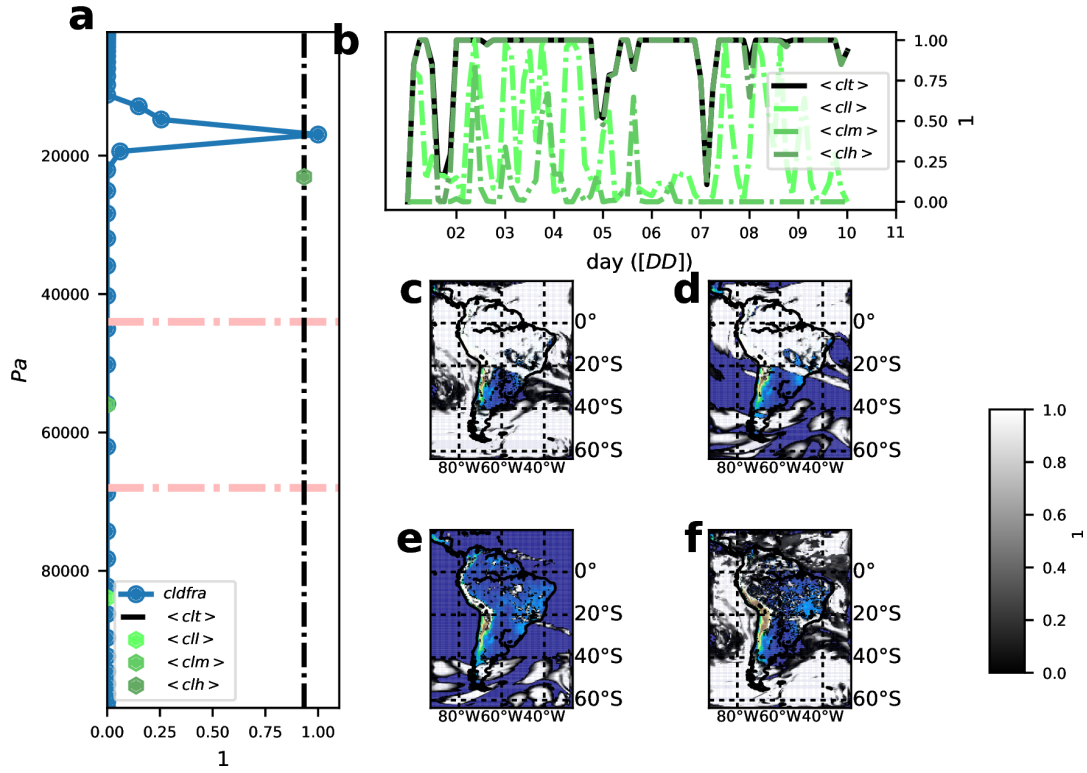


Figure 4. Vertical distribution of cloud fraction and the different cloud layers on 2012-12-09 at 15 UTC at $S 62^{\circ} 4' 38.00''$, $4^{\circ} 58' 55.51'' W$ (a): cloud fraction (*cldfra*, full circles with line in blue), mean total cloud fraction (*clt*, vertical dashed line), mean low-level cloud fraction (*ccl* $p \geq 680$ hPa, dark green hexagon), mean mid-level (*clm* $680 < p \geq 440$ hPa, green hexagon), mean high-level (*clh* $p < 440$ hPa, clear green hexagon). Temporal evolution of cloud layers at the given point (b). Map of *clt* with colored topography beneath to show-up cloud extent (c), map of *clh* (d), map of *clm* (e) and map of *ccl* (f)

$$\begin{aligned}
 & z_{clear} = 1, z_{cloud} = 0, ZEPSEC = 1.0 \times 10^{-12} \quad (15) \\
 & iz = 1 \text{ to } dimz \left\{ \begin{aligned}
 & z_{clear} = z_{clear} \frac{1 - MAX[CLDFRAC(iz), z_{cloud}]}{1 - MIN[z_{cloud}, 1 - ZEPSEC]} \\
 & clt = 1 - z_{clear} \\
 & z_{cloud} = CLDFRAC(iz)
 \end{aligned} \right.
 \end{aligned}$$

where *CLDFRAC*: cloud fraction (1) at each vertical level, *zclear*: clear-sky value (1), *zcloud*: cloud-sky value (1), *ZEPSEC*: value for very tiny number. Same methodology as in equation 15 is applied for the diagnostic of ‘clh’, ‘clm’ and ‘ccl’ but splitting by corresponding pressure layers. The figure 4 illustrates the result of the implementation and compare the results with the actual values of the cloud fraction (a and b panels) as well as the different cloud distribution (panels c to f).

Wind derived variables

CORDEX requires two wind-derived diagnostics: the daily maximum near-surface wind speed of gust (`wsgsmax`) and the daily maximum wind speed of gust at 100m above ground (`wsgsmax100`). These variables can not be retrieved by post-processing the standard output since they require the combination of different variables (some of them not available from the model output)

5 and have to be produced as a maximum value. The module provides different ways to compute them under certain limitations.

`wsgsmax`: maximum near-surface wind speed of gust

The wind gust accounts for the wind from upper levels that is projected to the surface due to instability within the planetary boundary layer. In the current version of the module two complementary methods of diagnosing the variable have been implemented (resultant winds are Earth-rotated). The choice between the two methods is done by the parameter labeled `wsgs_diag`

10 (in `cordex` section), with the default value set to 1. The implemented methods are:

- The Brasseur method [`wsgs_diag` = 1]: An implementation of wind gust considering Turbulent Kinetic Energy (TKE) estimates and stability defined by virtual temperature (θ_v) as indicated in equation 16 following Brasseur (2001). Implementation is adapted from a version already introduced in the CLimate WRF (`clWRF`, <http://www.meteo.unican.es/wiki/cordexwrf/SoftwareTools/ClWrf>; Fita et al., 2010),

$$15 \quad \frac{1}{z_p} \int_0^{z_p} TKE(z) dz \geq \int_0^{z_p} g \frac{\Delta\theta_v(z)}{\Theta_v(z)} dz \quad (16)$$

where TKE : Turbulent Kinetic Energy (m^2s^{-2}), θ_v : virtual potential temperature (K). z_p height of the considered parcel (m , maximum height which satisfies equation 16), $\Delta\theta_v(z)$: variation of θ_v over a given layer (Km^{-1}).

- The AFWA method [`wsgs_diag` = 2]: An implementation adopted from the WRF module `module_diag_afwa.F` which calculates the wind gust that only occurs as a blending of upper-level winds $zagl$ (around 1 km above ground; $zagl(k_{1000}) \geq 1000 m$, see equation 17) when precipitation intensity per hour is above a given maximum value, $prate_{hr}^{mm} \geq 50 mmh^{-1}$

$$20 \quad \begin{aligned} \mathbf{va}_{1km} &= \mathbf{va}(k_{1000} - 1) + [1000 - zagl(k_{1000} - 1)] \frac{\mathbf{va}(k_{1000}) - \mathbf{va}(k_{1000} - 1)}{zagl(k_{1000})} \\ \gamma &= \frac{150 - prate_{hr}^{mm}}{100} \\ \mathbf{va}_{blend} &= \mathbf{vas}\gamma + \mathbf{va}_{1km} \times (1 - \gamma) \end{aligned} \quad (17)$$

25 where va : air wind (ms^{-1}), $zagl$: height above ground (m), k_{1000} : vertical level at which $zagl$ is equal or above 1000 m . $prate_{hr}^{mm}$: hourly precipitation rate (mmh^{-1}), va_{blend} : blended wind (ms^{-1})

The two methods provide wind gust estimation (WGE) from two different perspectives: mechanic and convective. In order to take into account both perspectives, additional variables: `totwsgsmax` (total maximum wind-gust speed at surface), `totugsmax` (total maximum wind-gust eastward speed at surface), `totvgmax` (total maximum wind-gust northward speed at

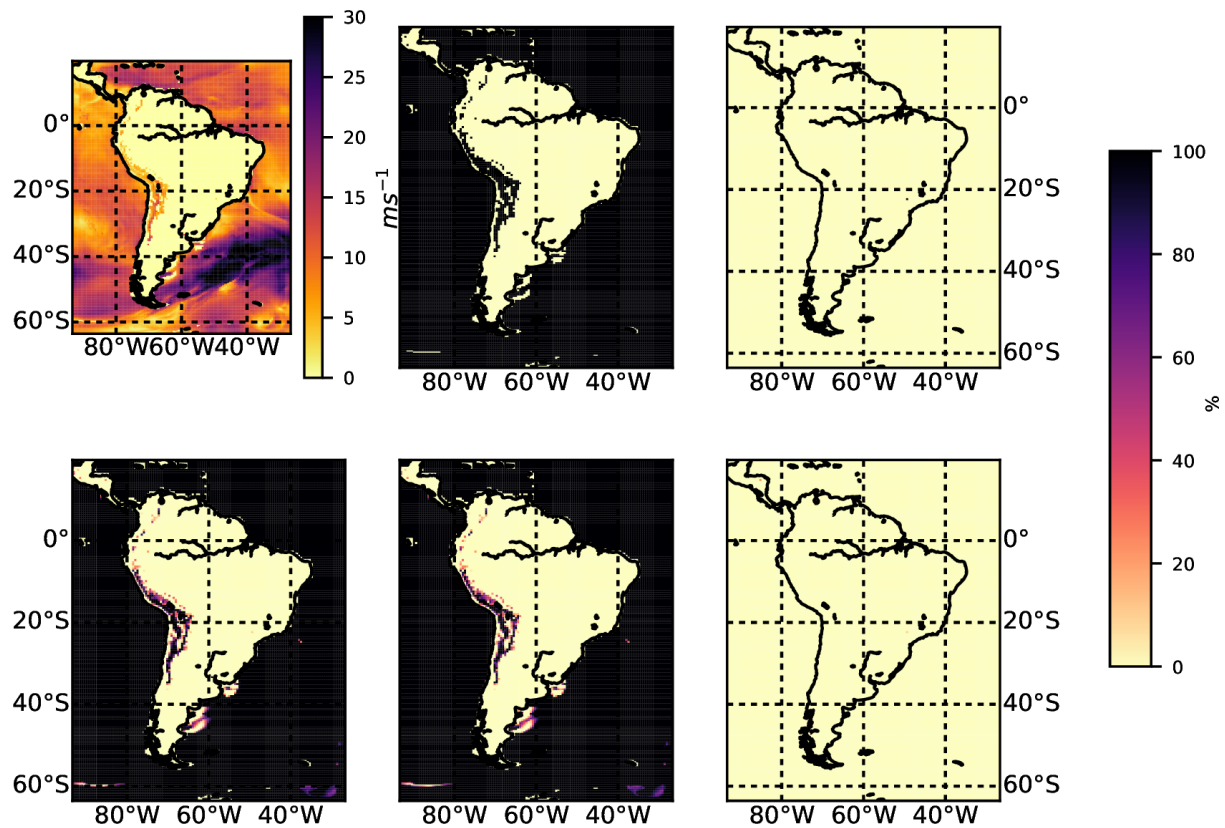


Figure 5. Near surface wind gust estimates on 2012 December the 9th at 15 UTC. 3h-maximum total wind gust strength ($wsgsmax^{tot}$, top left), percentage of $wsgsmax^{tot}$ following the Brasseur method ($wsgsmax^{b01}$, top middle), percentage following the AFWA-heavy precipitation implementation ($wsgsmax^{hp}$, top right), percentage of time steps where grid point got total wind gust (bottom left), percentage of time steps where grid point got $wsgsmax^{b01}$ (bottom middle), percentage due to $wsgsmax^{hp}$ (bottom right)

surface), and $totwsgspercen$ (percentage of time steps along 9freq in which grid point got wind gust in %). Figure 5 shows the outcomes when applying each method. It is shown how wind gust is mainly originated by turbulence, with a minor impact of heavy precipitation events at the depicted time. Furthermore, in the bottom panel it is shown how wind-gusts are highly frequent above the sea in comparison to the low frequency above continental flat areas (Andes mountain range exhibits high occurrence of wind gust).

wsgsmax100: daily maximum wind speed of gust at 100 m The calculation of wind gust at 100 m should follow a similar implementation used for calculating the $wsgsmax$, but at 100 m. An extrapolation of such turbulent phenomena would require a complete new set of equations which have not been placed yet. However, it could be considered as first approach to take the same wind gust as the one at the surface (when it is deflected from above 100 m). The assumption would be that the wind gust at 100 m would correspond to the deflected wind on its ‘way’ to the surface. Instead, as a way to complement, the estimation

of maximum wind speed at 100 m is provided. Provided wind components are also Earth-rotated. Three different methods have been implemented. Two following per-assumed vertical wind profiles (after PhD thesis Jourdir, 2015) and a third one following Monin-Obukhov theory to estimate the wind components above ground. These three methods are chosen by the namelist parameter labeled `wsz100_diag`. Its default value is 1. The implemented methods are:

- 5 – [`wsz100_diag` = 1], following the power-law wind vertical distribution as depicted in equation 18 using the upper-level atmospheric wind speed below ($k_{100}^<$) and above ($k_{100}^>$) the height above ground of 100 m (z_{agl}),

$$\begin{aligned} \mathbf{va}_{100} &= \mathbf{va}(k_{100}^>) \left(\frac{100.}{z_{agl}(k_{100}^>)} \right)^{\alpha_{x,y}} \\ \alpha_{x,y} &= \frac{\ln(\mathbf{va}(k_{100}^>)) - \ln(\mathbf{va}(k_{100}^<))}{\ln(z_{agl}(k_{100}^>)) - \ln(z_{agl}(k_{100}^<))} \end{aligned} \quad (18)$$

- 10 – [`wsz100_diag` = 2], following logarithmic-law wind vertical distribution, as it is depicted in equation 19, using upper-level atmospheric wind speed below ($k_{100}^<$) and above ($k_{100}^>$) the height above ground of 100 m (z_{agl})

$$\begin{aligned} \ln(z_0) &= \frac{\mathbf{va}(k_{100}^>) \ln(z_{agl}(k_{100}^<)) - \mathbf{va}(k_{100}^<) \ln(z_{agl}(k_{100}^>))}{\mathbf{va}(k_{100}^>) - \mathbf{va}(k_{100}^<)} \\ \mathbf{va}_{100} &= \mathbf{va}(k_{100}^>) \frac{\ln(100.) - \ln(z_0)}{\ln(z_{agl}(k_{100}^>)) - \ln(z_0)} \end{aligned} \quad (19)$$

- 15 – [`wsz100_diag` = 3], following Monin-Obukhov theory. User should keep in mind that this method is not useful for heights larger than few decimeters ($z > 80$ m). The wind at given height is extrapolated following turbulent mechanisms. As it is shown in equation 20, surface wind speed is used as surrogate to estimate 100 m wind direction ($\theta_{10} = \tan^{-1}(uas, vas)$, without considering Eckman pumping, or other effects on wind direction). In this implementation u^* in similarity theory is taken as WRF estimates UST, Monin-Obukhov length (\mathcal{L}_O) as the WRF values RMOL, roughness length (z_0) as WRF thermal time-varying roughness length ZNT,

$$\begin{aligned} wss_{100} &= \frac{UST}{\kappa} \left(\ln \left(\frac{100}{z_0} \right) + \Psi_M \left(\frac{100}{\mathcal{L}_O} \right) \right) \\ \mathcal{L}_O &= \frac{-UST^3 T_v}{\kappa g Q_0} \quad (Obukhov \text{ length}) \end{aligned} \quad (20)$$

$$\Psi_M \left(\frac{z}{\mathcal{L}_O} \right) = \begin{cases} \frac{4.7z}{\mathcal{L}_O} & \frac{z}{\mathcal{L}_O} > 0 \quad (sTable) \\ \ln \left[\left(\frac{1+X^2}{2} \right) \left(\frac{1+X}{2} \right)^2 \right] - 2 \tan^{-1}(X) + \frac{\pi}{2} & \frac{z}{\mathcal{L}_O} < 0 \quad (unsTable) \end{cases}$$

$$X = \left(1 - \frac{15z}{\mathcal{L}_O} \right)^{1/4}$$

$$\theta_{10} = \text{atan} \left(\frac{V10}{U10} \right) \mathbf{va}_{100} = \begin{cases} ua_{100} = wss_{100} \cos(\theta_{10}) \\ va_{100} = wss_{100} \sin(\theta_{10}) \end{cases}$$

- 25 wss_{100} : wind speed at 100 m (ms^{-1}), where Ψ_M : stability function after (Businger et al., 1971), UST : u^* in similarity theory (ms^{-1}), z_0 : roughness length (m), $U10, V10$: 10-m wind speed (ms^{-1}), θ_{10} : 10-m wind speed direction (rad),

Table 3. Mixing ratio associated with column integrated variables

name	WRF species	description
prw	QVAPOR	water vapor mixing ratio
clwvi	QCLOUD+QRAIN	condensed water and rain mixing ratio
clivi	QICE+QSNOW+QGRAUPEL+QHAIL	ice, snow, graupel and hail mixing ratio
clgvi	QGRAUPEL	graupel mixing ratio
clhvi	QHAIL	hail mixing ratio

ua_{100} : 100 m eastward wind speed (ms^{-1}), va_{100} : 100 m northward wind speed (ms^{-1}). Note the absence of correction in wind direction to Ekman pumping or other turbulence effects.

The user can also select the height at which the estimation is computed throughout the namelist parameter `z100m_wind` (100 m as default value). Figure 6 shows different preliminary results using the three different approximations. It is illustrated (a panel) how wind-gusts are larger than the 10-m diagnostic winds, and also the difference is larger when using Monin-Obukhov method compared to the two others methods. Certain problems (too small Monin-Obukhov length) are recognized when applying Monin-Obukhov for extrapolating wind at 100 m, which is shown in panel b, where wind gusts appear to be strong as $80 ms^{-1}$. Therefore a user is advised to use this method with care.

Vertically integrated variables

- 10 The instantaneous vertically integrated amount of water vapor (prw), liquid condensed water species (clwvi), ice species (clivi), graupel (clgvi), hail (clhvi) are the vertically integrated amounts of each species along the vertical column (density weighted) over each grid point. They are provided using the same implementation as those in `p_interp.F` - WRF tool for vertical interpolation. The general equation following WRF standard variables is:

$$clvivar = \frac{MU + MUB}{g} \sum_{iz=1}^{e_vert} WRFVAR[iz](DNW[iz]) \quad (21)$$

- 15 where *clvivar*: the column vertically integrated variable's CF-compliant name (prw, clwvi, clivi, clgvi, or clhvi), *MU*: perturbation dry air mass in column (*Pa*), *MUB*: base-state dry air mass in column (*Pa*), *g*: gravity (ms^{-2}), *e_vert*: total number of vertical levels, *WRFVAR*: the water species' mixing ratio at each sigma level ($kgkg^{-1}$), *DNW*: difference between two consecutive full-eta levels (-). Table 3 indicates the *WRFVAR* names associated with the *clvivar* names.

- Note that clgvi and clhvi are part of the 'Tier1' level and are only accessible if pre-compilation variable `CDXWRF` is set to 1. See section 2.3 for more detail. In order to provide an example of the correct computation of the diagnostics, results at a given grid point are shown in figure 7. It is shown that the total precipitable water (prw) correctly corresponds to the density weighted vertical integration of the water content along the column of air.

evspsblpot: potential evapotranspiration

This variable represents the evaporative demand of the atmosphere. It is computed following the standard method already

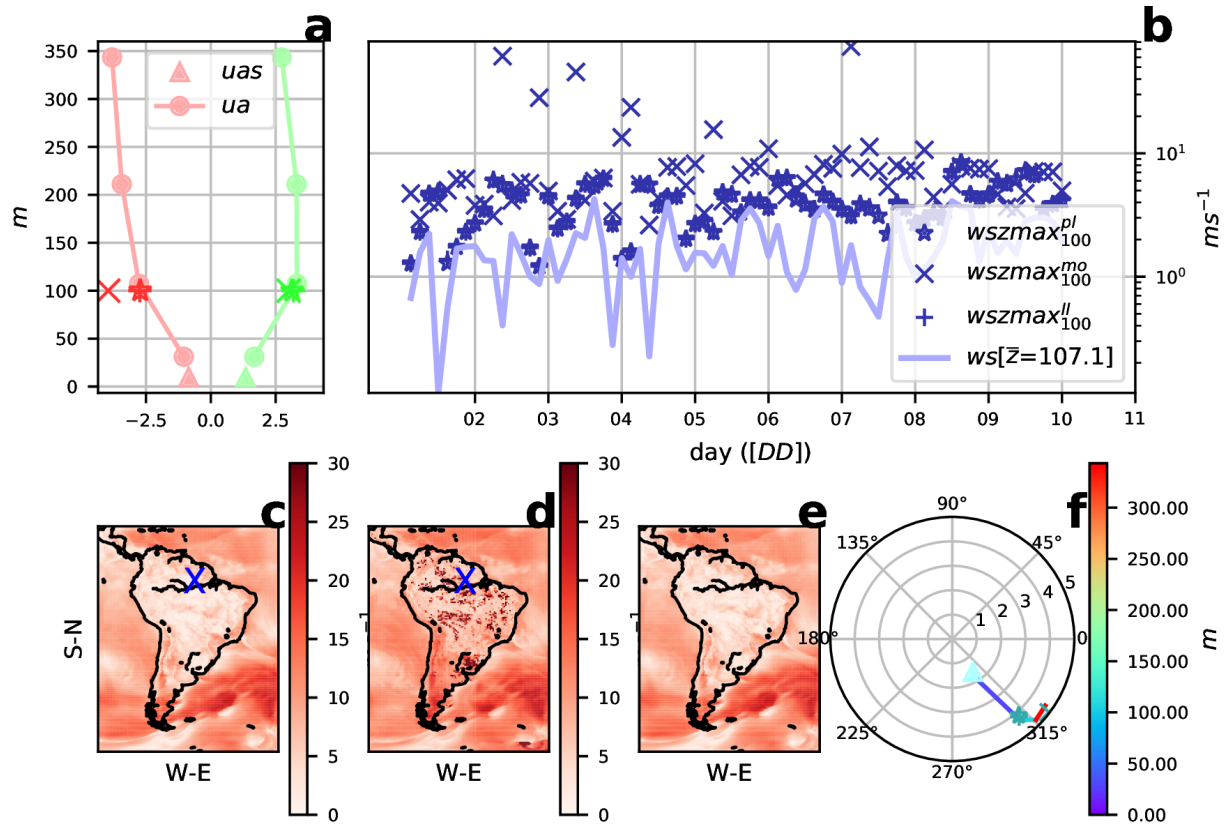


Figure 6. 100 m wind estimates. Comparison between upper-level winds and estimation on 2012-12-09 at 15 UTC at $S 62^{\circ} 4' 38.00''$, $4^{\circ} 58' 55.51'' W$ (a): 3h-maximum eastward wind (red) at 100 m by power-law ($uzmax^{pl}$, star marker), Monin-Obukhov theory ($uzmax^{mo}$, cross) by logarithmic law ($uzmax^{ll}$, sum) 10-m wind value (uas , filled triangle) and upper-level winds (ua , filled circles with line), also for the northward component (green). Temporal evolution of wind speed (b) with all approximations and upper-level winds at the closest vertical level at 100 m (on log-y scale, $\bar{z} = 107.1$ m on average). Maps of three estimations: power-law (c), Monin-Obukhov (d), logarithmic-law (e) with the blue cross showing the point of previous figures. Vertical distribution of winds at the given point in Wind rose-like representation (f)

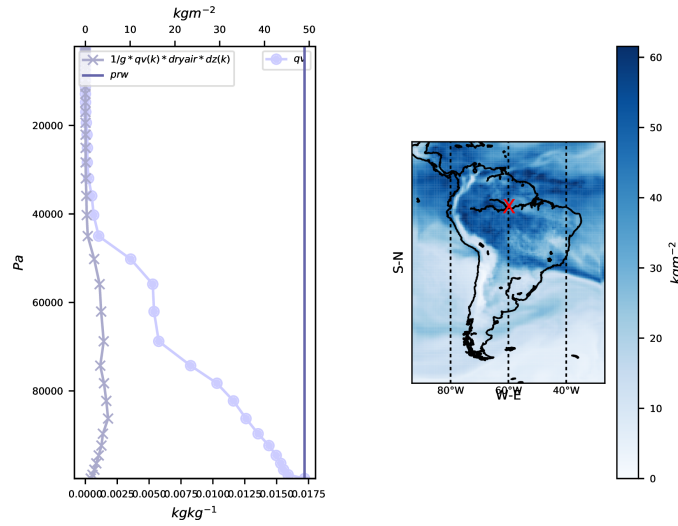


Figure 7. On 2012-12-09 at 15 UTC at $S 62^{\circ} 4' 38.00''$, $4^{\circ} 58' 55.51'' W$ (left): water path (prw , vertical straight line in mm top x-axis), vertical profile of water vapour (qv , line with full circles in $kgkg^{-1}$ bottom x-axis), water path at each level (line with crosses). Map of water path (right) on 2012-12-09 at 15 UTC, red cross shows where the vertical accumulation is retrieved

implemented in most GCMs. One of the first proposed methods was provided by Manabe (1969). Some corrections have been proposed to the initial methodology in order to overcome its deficiencies (e.g. see Barella-Ortiz et al. (2013) for an intercomparison among different methods). It is provided as an averaged flux. Calculation of the potential evapotranspiration can be activated with the namelist input parameter `potevap_diag` (number 2 is the default option):

- 5 – bulk method [`potevap_diag = 1`]: this method corresponds to the original one proposed in Manabe (1969). It basically consists in a difference between a supposed saturated air at the surface temperature and the humidity of the atmosphere as it is depicted in equation 22,

$$qc = C_D \sqrt{U10^2 + V10^2}$$

$$evspsblpot_{bulk} = \rho(1)qc[ws(ts) - QVAPOR(1)] \quad (22)$$

- 10 where $ws(ts)$: saturated air at ts ($kgkg^{-1}$), qc : surface drag coefficient (ms^{-1}), TSK : surface temperature (K), $ws(ts)$: saturated air by surface temperature ($kgkg^{-1}$) based on August-Roche-Magnus approximation, $press$: air pressure (Pa), $U10, V10$: 10 m wind components (ms^{-1}), $QVAPOR$: 3D water vapour mixing ratio ($kgkg^{-1}$), C_D : drag coefficient (–, only available from MM5-similarity and MYNN surface layer schemes, otherwise is zero).

- 15 – Milly92 method [`potevap_diag = 2`]: this method makes a correction of the bulk diagnostic by introducing a Milly's correction parameter ξ , which accounts for other atmospheric-related phenomena (Milly, 1992). It is explained in equation 23, and its implementation is similar to the one present in ORCHIDEE model (Organising Carbon and

Hydrology In Dynamic Ecosystems, <http://orchidee.ipsl.fr/>, de Rosnay et al. (2002). The implementation is retrieved from the module `src_sechiba/enerbil.f90`,

$$\begin{aligned}
 \beta &= \frac{sfcevap}{evspsblpot_{bulk}} \\
 \partial_T ws(T) &= \frac{ws[T(1) + 0.5] - ws[T(1) - 0.5]}{2 \times 0.5} \\
 \xi &= \frac{L\rho(1)qc\partial_T ws(T)(1 - \beta)}{4EMISSCtBoltzmanT(1)^3 + \rho(1)Cpqc + L\rho(1)qc\partial_T ws(T)\beta} \\
 evspsblpot_{Milly92} &= evspsblpot_{bulk} \frac{1}{1 + \xi}
 \end{aligned} \tag{23}$$

where β : Moisture availability function, $sfcevap = QFX$ surface evaporation ($kgm^{-2}s^{-1}$) from QFX : surface moisture flux ($kgm^{-2}s^{-1}$), L : latent heat of vaporization, $EMISS$: emissivity (1), $CtBoltzman$: constant of Stefan-Boltzman, Cp : specific heat of air, $\partial_T ws(T)$: derivative of saturated air at temperature of the first atmospheric layer ($kgkg^{-1}$) using numerical 1st order approximation

See in figure 8 an example of the differences between both implementations. It shows how important is the correction introduced by Milly and its strong effect on the diurnal cycle. Basically, the correction permits potential evapotranspiration during night time and reduces its strength at noon (18 UTC corresponds to 12 local time). There is a generic diagnostic to overcome boundary layer scheme dependency in the calculation of the drag coefficient (see below in generic variables).

3.3 Generic variables

Some of the diagnostics required by CORDEX depend on the approximations, equations, methodologies and observations used to compute them. This makes model intercomparison very difficult, since values might differ from one implementation to another. In order to overcome this problem, a series of variables are also provided in a ‘generic’ form (when possible) meaning that they are obtained directly from standard variables. Thus, these generic forms of the diagnostics become ‘independent’ of the model’s implementation.

edgen: generic surface drag coefficient

Computation of the instantaneous drag coefficient at the surface depends on the selected surface scheme. In order to avoid this scheme dependency, a generic calculation of the coefficient has been introduced as in equation 24 following Garratt (1992),

$$C_D^{gen} = \left(\frac{UST}{wss} \right)^2 \tag{24}$$

with UST : u^* being friction velocity from the similarity theory (ms^{-1}), and $wss = \sqrt{U10^2 + V10^2}$: 10-m wind speed (ms^{-1})

tauvgen: generic surface downward wind stress

Generic surface downward wind stress at 10m is calculated following the equation 25 which uses the generic diagnostics of the drag coefficient.

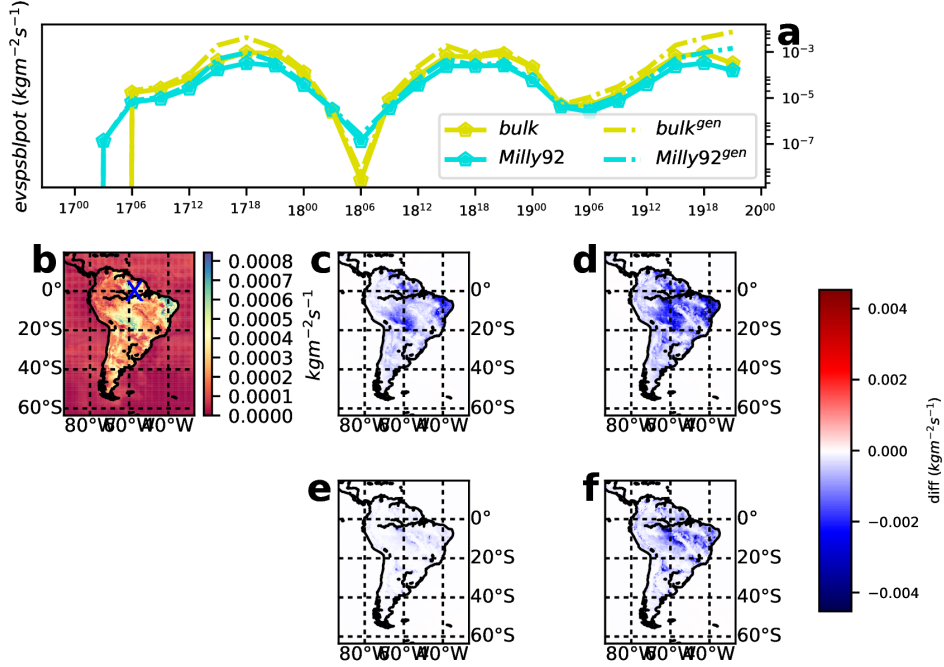


Figure 8. Evolution (a, in y-log scale) of potential evapotranspiration by bulk (yellow) and Milly92 (blue) generic (dashed lines) methods at $S 4^{\circ} 58' 55.524''$, $62^{\circ} 4' 37.92'' W$ (blue cross in b). On 2015-11-18 15 UTC, potential evapotranspiration following Milly correction (b), differences between both methods (c, $evspblpot_{Milly92} - evspblpot_{bulk}$), differences between both generic methods (d, $evspblpot_{Milly92}^{gen} - evspblpot_{bulk}^{gen}$), differences between Milly method and its generic counterpart (e, $evspblpot_{Milly92} - evspblpot_{Milly92}^{gen}$) and differences between bulk method and its generic counterpart (f, $evspblpot_{bulk} - evspblpot_{bulk}^{gen}$)

$$tauwgen = (C_D^{gen} uas^2, C_D^{gen} vas^2) \quad (25)$$

where C_D^{gen} : generic drag coefficient (–, see equation 24), uas : Earth-rotated eastward 10 m wind component, vas : Earth-rotated northward 10 m wind component.

rsusgen: surface upwelling shortwave radiation

- Surface upwelling shortwave radiation is the shortwave radiation directed from the surface. It is calculated in a generic way as the reflected shortwave radiation depending on the surface albedo as it is shown in equation 26,

$$rsus = -ALBEDO * SWDOWN \quad (26)$$

Being, *ALBEDO*: albedo (1), *SWDOWN*: downward at surface shortwave radiation (Wm^{-2})

rlusgen: surface upwelling longwave radiation

Surface upwelling longwave radiation is the longwave radiation coming from the surface. It is calculated in a generic way as the longwave radiation from a black body due to surface temperature following the Stefan-Boltzmann law as it is given in
5 equation 27,

$$rlus = CtBoltzman * EMISS * TSK^4 \quad (27)$$

Being, *CtBoltzman*: the Stefan-Boltzmann constant ($= 5.67051E^{-8} Wm^{-2}K^{-4}$), *EMISS*: surface emissivity (1), *TSK*: skin temperature (K)

evspsblpotgen: generic potential evapotranspiration

10 This variable corresponds to the generic definition of potential evapotranspiration ('evspsblpot'). The same two methodologies as in the the regular diagnostic have been implemented. The only difference is that in this case, the generic estimation of the drag coefficient 'cdgen' is used (see equation 24) instead of the one given by the model.

3.4 Tier1 variables

These diagnostics are required by CORDEX, but they are not mandatory. They have been also included as a way to fulfill all
15 the CORDEX requirements. These variables require the setting of the pre-compilation flag `CDXWRF` to 1 and performing some complementary modifications in the module's Registry file `registry.cordex`. See section 2.3 for more details.

zmlagen: generic boundary layer height

Instantaneous Planetary Boundary Layer (PBL) height is a clear example of model dependence and even scheme dependence of how a diagnostic is computed. Each PBL scheme has its own assumptions and 'zmla' is computed in a scheme-dependent
20 specific way.

In order to overcome the model/scheme dependence, we implemented a generic formulation for calculating the PBL height as it was done in (García-Díez et al., 2013) after (Nielsen-Gammon et al., 2008). The method consists in defining the height of the PBL as the first level in the mixed layer (ML) where potential temperature exceeds the minimum potential ML temperature by more than 1.5 K . It has been implemented using the definitions given below:

- 25 1. Mixed layer depth (MLD) is defined as the model level (k_{MLD}) starting from the second model level at which the variation of the mixing ratio ($QVAPOR(k)$, normalized with its value at the first level) exceeds some predefined threshold value ($QVAPOR(1)$): $\frac{|QVAPOR(k_{MLD}) - QVAPOR(1)|}{QVAPOR(1)} > \delta QVAPOR$ (here applied a $\delta QVAPOR = 0.1$)
2. Within the MLD the value with the minimum potential temperature is taken as: $\theta_{min_{MLD}} = \min[\theta(1), \dots, \theta(k_{MLD})]$
3. The level of the PBL height (k_{zmla}) is the level at which the maximum variation of potential temperature within the
30 MLD exceeds some predefined threshold value: $\theta(k_{zmla}) - \theta_{min_{MLD}} > \delta\theta$, (here $\delta\theta = 1.5 K$)
4. The PBL height ($zmla$) is obtained using the geopotential height zg at the calculated k_{zmla} level above the ground ($zagl$): $zmla = zagl(k_{zmla}) = zg(k_{zmla})/g - HGT$, with HGT being surface elevation height above sea level.

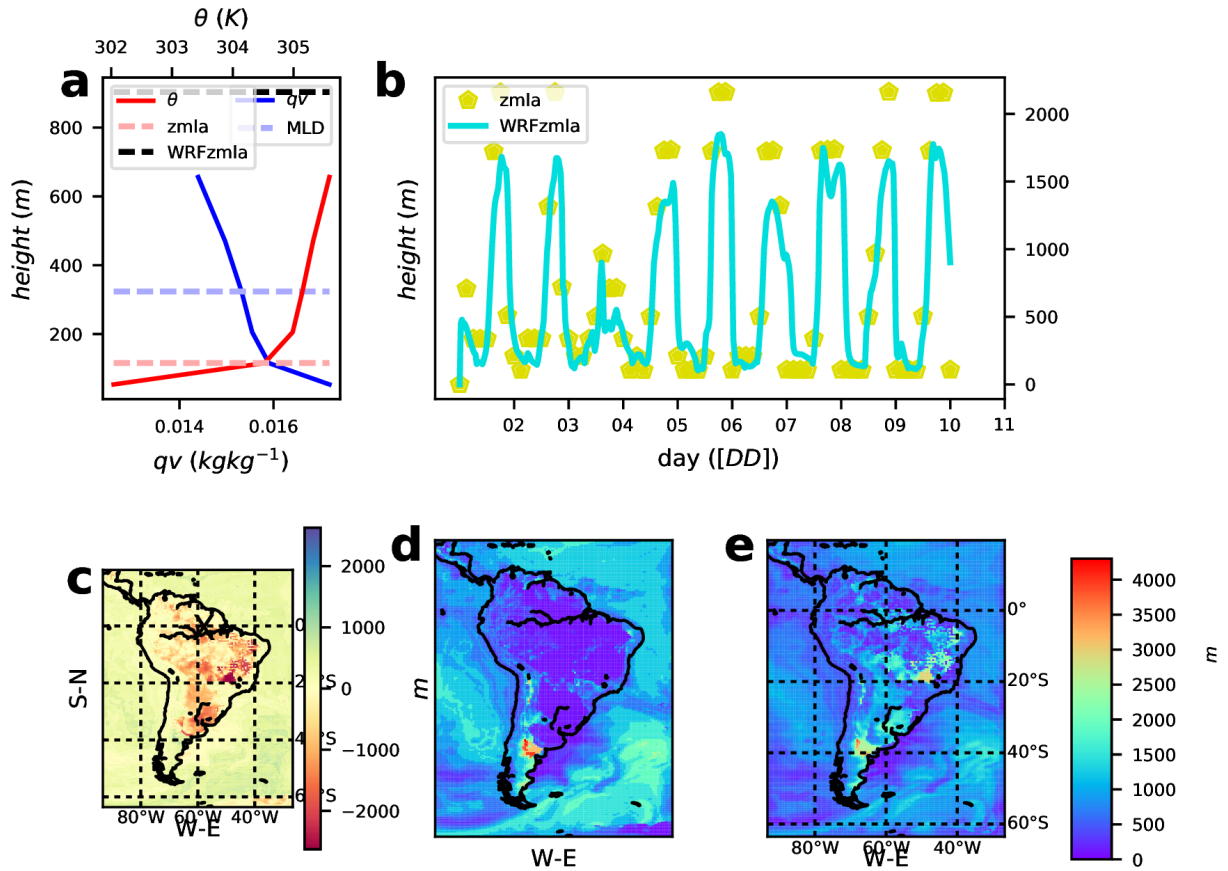


Figure 9. Vertical characteristics of the atmosphere on 2012-12-09 at 15 UTC at $S 62^{\circ} 4' 38.00''$, $4^{\circ} 58' 55.51'' W$ (a): potential temperature vertical profile (θ K, red line), vertical profile of mixing ratio (qv $kgkg^{-1}$, blue line), mixed layer depth (MLD , dashed horizontal line at 323.522 m), derived boundary layer height ($zmla$, horizontal dashed line at 1007.122 m) and WRF derived pbl scheme value ($WRFzmla$ at 903.017 m). Comparison of temporal evolutions (b) between derived $zmla$ (yellow stars) and WRF's pbl scheme (blue line). Map of differences between derived and WRF simulated ($zmla - zmla_{WRF}$, c), $zmla$ map (d) and $zmla_{WRF}$ (e)

No general rule has been applied to provide the correct value of δqv used to determine MLD . It can be determined by the namelist parameters $zmlagen_dqv$ for δqv (default value 0.1) and $zmlagen_dtheta$ for $\delta\theta$ (default value 1.5 K). Comparison of this implementation with the $zmla$ directly provided by WRF's Mellor-Yamada Nakanishi and Niino Level 2.5 PBL scheme (MYNN2.5 Nakanishi and Niino, 2006) is shown in figure 9. In general the generic estimation produces a higher PBL (a panel) with lower values during night (b panel). Spatial distributions between both diagnostics are pretty similar.

Convective diagnostics

Diagnostics related to convective activity are: Convective Available Potential Energy (CAPE) which accounts for all the energy that might be released convectively, Convective Inhibition (CIN) which accounts for processes which inhibit the convection,

Height of the Level of Free Convection (ZLFC), Pressure at the Level of Free Convection (PLFC), and Lifted Index (LI) which accounts for the temperature difference between the environmental temperature at some higher level in the troposphere and the temperature that a parcel would have if adiabatically lifted at that level. CORDEX requires these values as statistics between output times (9freq in this case)

- 5 Since the version V3.6 of WRF, these variables can already be calculated with the module `module_diag_afwa.F` via the `Buoyancy` function. In this version of the module, this is the only available implementation. These vertically integrated diagnostics have a high computational cost. In order to minimize it, they are only computed at output time step by default. However, if a user requires them as statistics (such as `capemin`, `capemax`, `capemean`), then these diagnostics are computed at all time steps. This behavior of the module is regulated via the namelist parameter `convxtrm_diag` (default value is
- 10 0, meaning no computation), and by setting the pre-compilation flag `CDXWRF` to 1 and performing some complementary modifications in module's Registry file `registry.cordex`. See section 2.3 for more detail.

4 Additional variables

- Some variables not required by CORDEX but which may be interesting and useful to the community for wide variety of the purposes have also been added. These variables will be obtained if the pre-compilation flag `CDXWRF` is set to 2 and some
- 15 additional modifications are made in module's registry file `registry.cordex`. See section 2.3 for more details.

tds: dew point temperature

The dew point temperature (cooler temperature at which air would saturate due to its current moisture content) is calculated following the August-Roche-Magnus approximation as it is shown in equation 28,

$$\gamma = \log(hurs) + \frac{b(T2 - 273.15)}{(T2 - 273.15) + c} \quad (28)$$

$$20 \text{ tds} = \frac{c\gamma}{b - \gamma} + 273.15$$

where $T2$: 2m temperature (K), $hurs$: 2m relative humidity (%), previously computed), $b = 17.625$, $c = 243.04$. This variable is provided as statistics: minimum, maximum and mean in the output.

Atmospheric water budget

- The water budget accounts for all the dynamics of the water in the atmosphere. This budget is divided in different terms
- 25 (dynamical and source/sink) accounting for the total mass of water. It can be computed independently for each water species. The equation for any given water species is given in equation 29:

$$\begin{aligned} TEN_q &= HOR_q + VER_q + MP_q \\ \frac{\partial q_q}{\partial t} &= -V_h \nabla q_q - w \frac{\partial q_q}{\partial z} + SO_q - SI_q \end{aligned} \quad (29)$$

- Where q stands for one of the six water species (vapor, snow, ice, rain, liquid, graupel, or hail) concentrations ($kgkg^{-1}$), V_h
- 30 stands for horizontal wind speed (ms^{-1}), w stands for the vertical wind speed (ms^{-1}) and MP for the loss or gain of water

due to cloud microphysical processes. The term in the left-hand side of the equation represents the water species tendency (TEN or 'PW'), referring to the difference between q at the model's previous time step and at the actual time step, divided by the time step. TEN equals to the horizontal advection (HOR or 'F', first term in right-hand side of the equation), the vertical advection (VER or 'Z', second term in right-hand side) and the sources (SO) or sink (SI) of atmospheric water due to MP .

- 5 All terms are expressed in $kgkg^{-1}s^{-1}$. However, SO and SI can not be provided because they are micro-physics dependent and make difficult to provide a generic formula for them.

In order to obtain the total column mass of water due to each term (in mm), an integration following eq. 30 is applied to each term of eq. 29 (similarly as in 21):

$$10 \quad -\frac{1}{g} \int_{P_{sfc}}^{P_{top}} \langle term \rangle dp \quad (30)$$

Following the methodology of Huang et al. (2014) and Yang et al. (2011); Fita and Flaounas (2018) implemented a new module in WRF in order to allow the computation of the water budget terms during model integration. This implementation is provided with the CORDEX module, but these variables are only provided as temporal accumulations (within 9freq) and vertical integrations in two forms: total column values and divided by the same layers as the cloud diagnostics (low, medium, 15 high). The accumulation of diabatic heating from the microphysics scheme is provided as a proxy of the sink/sources due to microphysics effects.

Preliminary results for all water species are shown in figures 10 and 11. Water vapour exhibits the largest values in both total tendency and horizontal advection. Dynamics of the other water species seems to be highly correlated with the presence of a storm system (lower right corner in the maps) or due to orthographic influences (existence of Andes mountain range can be 20 inferred).

Figures from 12 to 15 show temporal evolution and accumulated maps at a given time for all the water budget terms, decomposed for vapour (qv) and snow (qs). Accumulated maps are grouped into vertical levels as it is done with the clouds: $p \geq 68000 Pa$, $40000 \leq p < 68000 Pa$, $p < 40000 Pa$. Largest amounts of the budget terms are mainly found in low (high) 25 levels for water vapour (snow), temporal evolution at a given point show complexity of the water dynamics with the terms compensating each other. Figures is also shown how contribution to the total diabatic term is large at low levels over the ocean (showing the role of evaporation) and larger at high levels above the continent.

fogvis: visibility inside fog

Fog is one of major causes of transportation disruption. The horizontal resolutions of state-of-the-art CORDEX activities like FPS_Alps (3 km) open the possibility to explore phenomena such as fog which was impossible to be analyzed in previous 30 regional climate experiments. In order to be able to contribute in the analysis of fog phenomena, three different methods to calculate visibility have been introduced. Visibility is used to determine the presence of fog at a given moment. In order to provide a quantity with the density of the fog, only the visibility during a fog event is kept. The three methods are:

W.B. AC PW normalized by σ on 2012-12-09 15

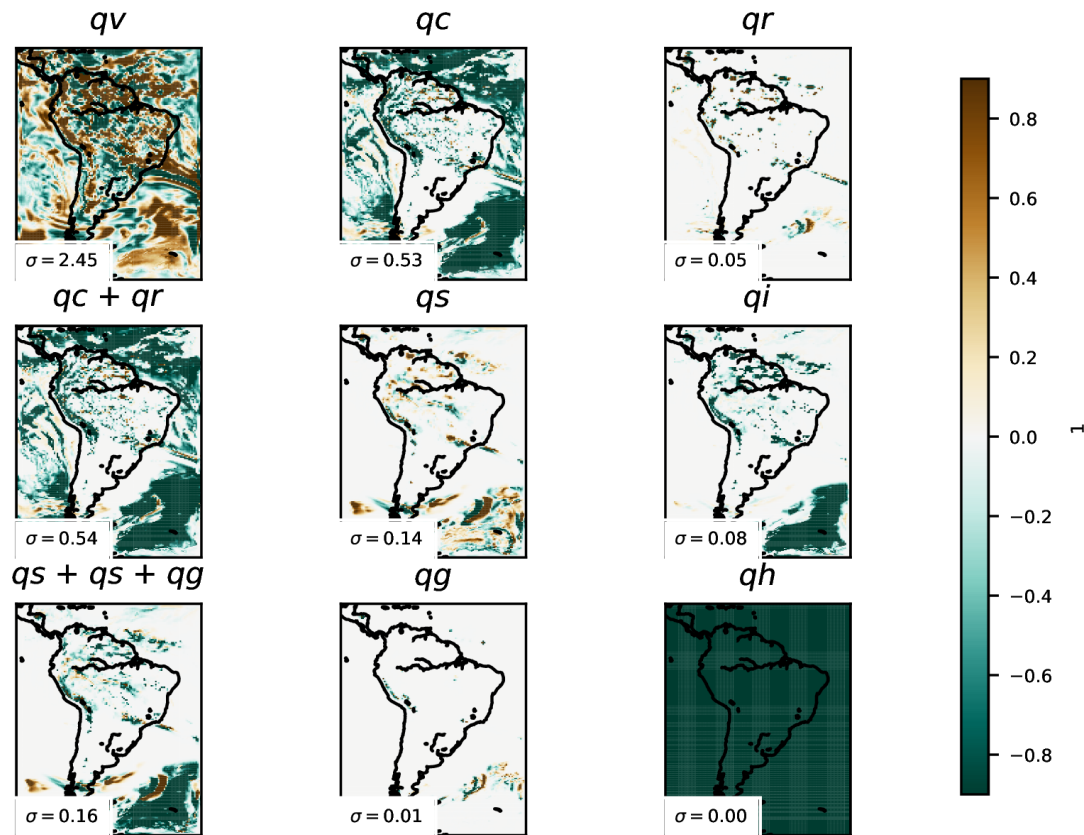


Figure 10. Normalized (with the spatial standard deviation of the mapped values, σ) water budget 3h-accumulated vertically integrated total tendency 'PW' at a given time, for water vapour (qv , top left), cloud (qc , top middle), rain (qr , top right), water condensed species ($qc + qr$, middle left), snow (qs , middle middle), ice (qi , middle right), water solid species ($qs + qi + qg$, bottom left), graupel (qg , bottom middle), hail (qh , bottom right). Number on low left corner of the figure correspond to the standard deviation (σ in mm) value used for the normalization

W.B. AC F normalized by σ on 2012-12-09 15

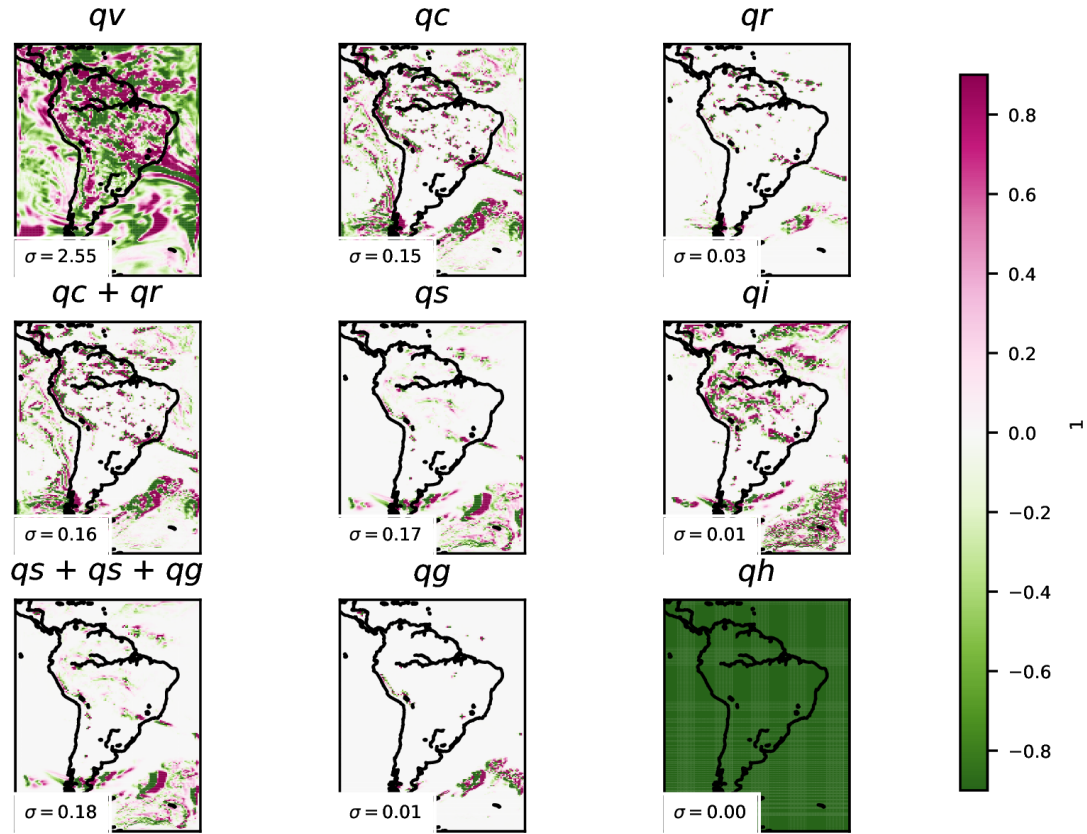


Figure 11. As in 10, but for Water budget 3h-accumulated vertically integrated horizontal advection 'F' at a given time

- **K84** [`visibility_diag = 1`]: Visibility is computed using liquid water (`Q_CLOUD`) and ice (`Q_ICE`) concentrations. Following (Bergot et al., 2007), fog appears when there are liquid and/or ice water species at the lowest model level present. Visibility is computed using equation 31 as in Kunkel (1984),

$$fogvis = \begin{cases} visc = 0.027(Q_{CLOUD} \times 1000)^{-0.88} & Q_{CLOUD} \neq 0 \\ visi = 0.024(Q_{ICE} \times 1000)^{-1.0} & Q_{ICE} \neq 0 \end{cases} \quad (31)$$

5 $fogvis = \min(visc, visi)$

where Q_{CLOUD} : liquid water (cloud) mixing ratio ($kgkg^{-1}$), Q_{ICE} : ice mixing ratio ($kgkg^{-1}$). Visibility values are in km

Water Budget terms at $-62^{\circ} 4' 38.00''$, $-4^{\circ} 58' 55.51''$ (60,150)

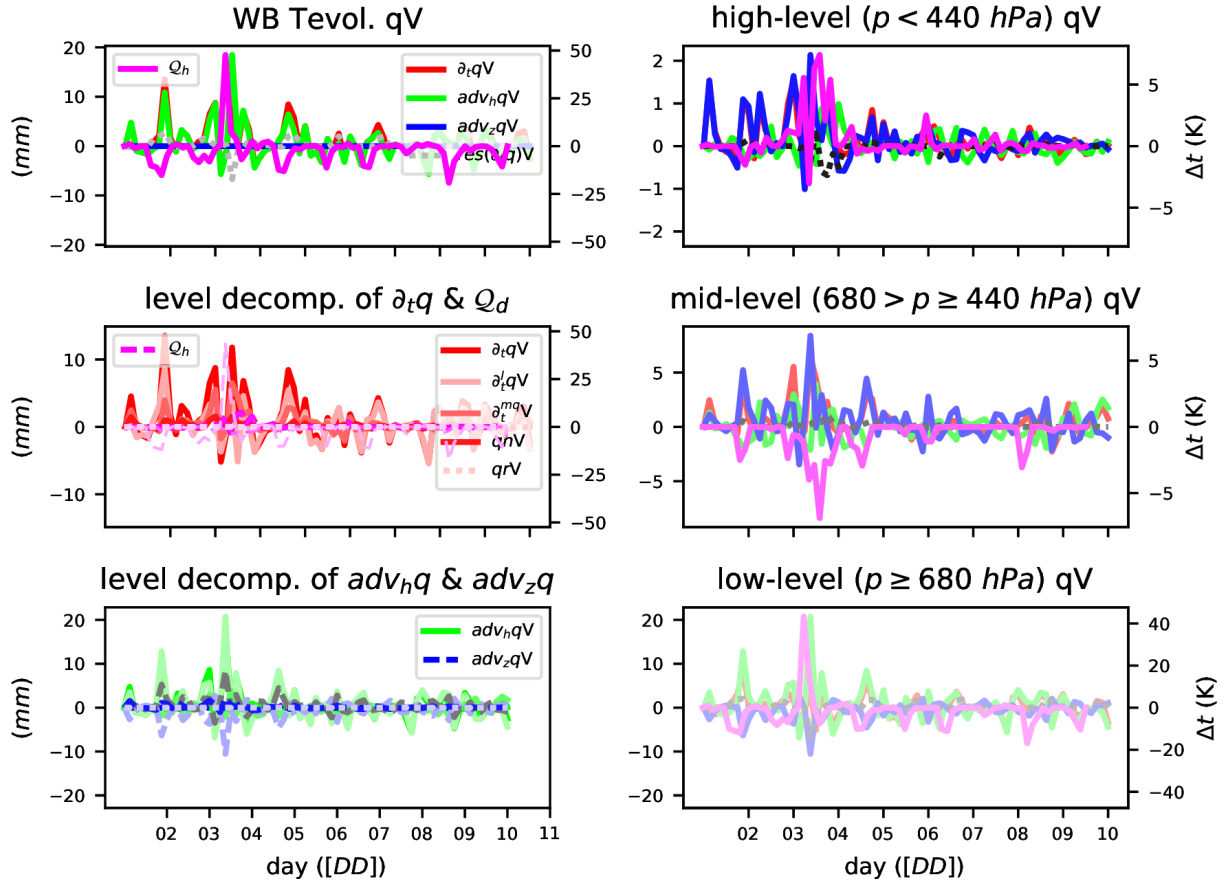


Figure 12. Water budget evolution at a given point for water vapour of vertically integrated water-budget terms: total tendency 'PW' ($\partial_t qv$, red), horizontal advection 'F' ($adv_h qv$, green), vertical advection 'Z' ($adv_z qv$, green), residual PW - F - Z ($res(\partial_t qv)$, gray dashed) and diabatic heating from micro-physics (Q_d , pink) (top left), only high-level vertically integrated values ($p < 440$ hPa, top right), high/mid/low-level (degree of color intensity) decomposition of $\partial_t qv$ (red) and Q_d (pink) and their respective residuals as dashed lines (middle left), only mid-level vertically integrated values ($680 > p \leq 440$ hPa, middle right), high/mid/low-level (degree of color intensity) decomposition of $adv_h qv$ (green) and $adv_z qv$ (blue) and their respective residuals as dashed lines (bottom left) and only low-level vertically integrated values ($p \geq 680$ hPa, bottom right)

Water Budget terms on 2012-12-09 15 UTC

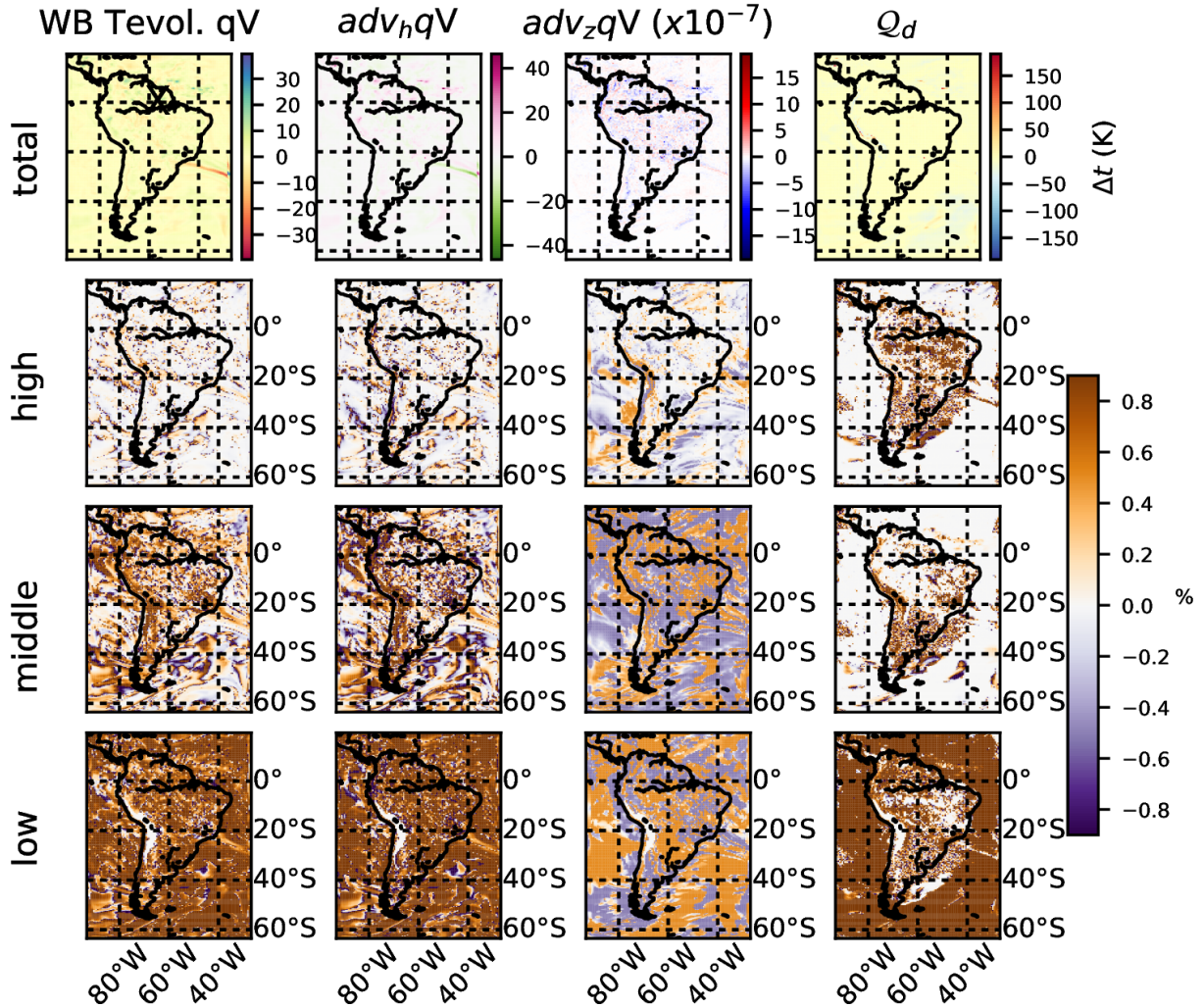


Figure 13. Water vapour budget maps of each component and diabatic heating from micro-physics at a given time and the percentual contribution at each different vertically integrated layer respective the total. total tendency 'PW' ($\partial_t qv$, first column), horizontal advection 'F' ($adv_h qv$, second col), vertical advection 'Z' ($adv_z qv$, third col.) and diabatic heating from micro-physics (Q_d , 4th col). Percentage contribution of high level ($p < 440 \text{ hPa}$) integration to the total (second row), percentage for mid level ($680 > p \geq 440 \text{ hPa}$) integration to the total (third row) and percentage of low-level ($p \geq 680 \text{ hPa}$) integration (bottom row)

Water Budget terms at -62° 4' 38.00", -4° 58' 55.51" (60,150)

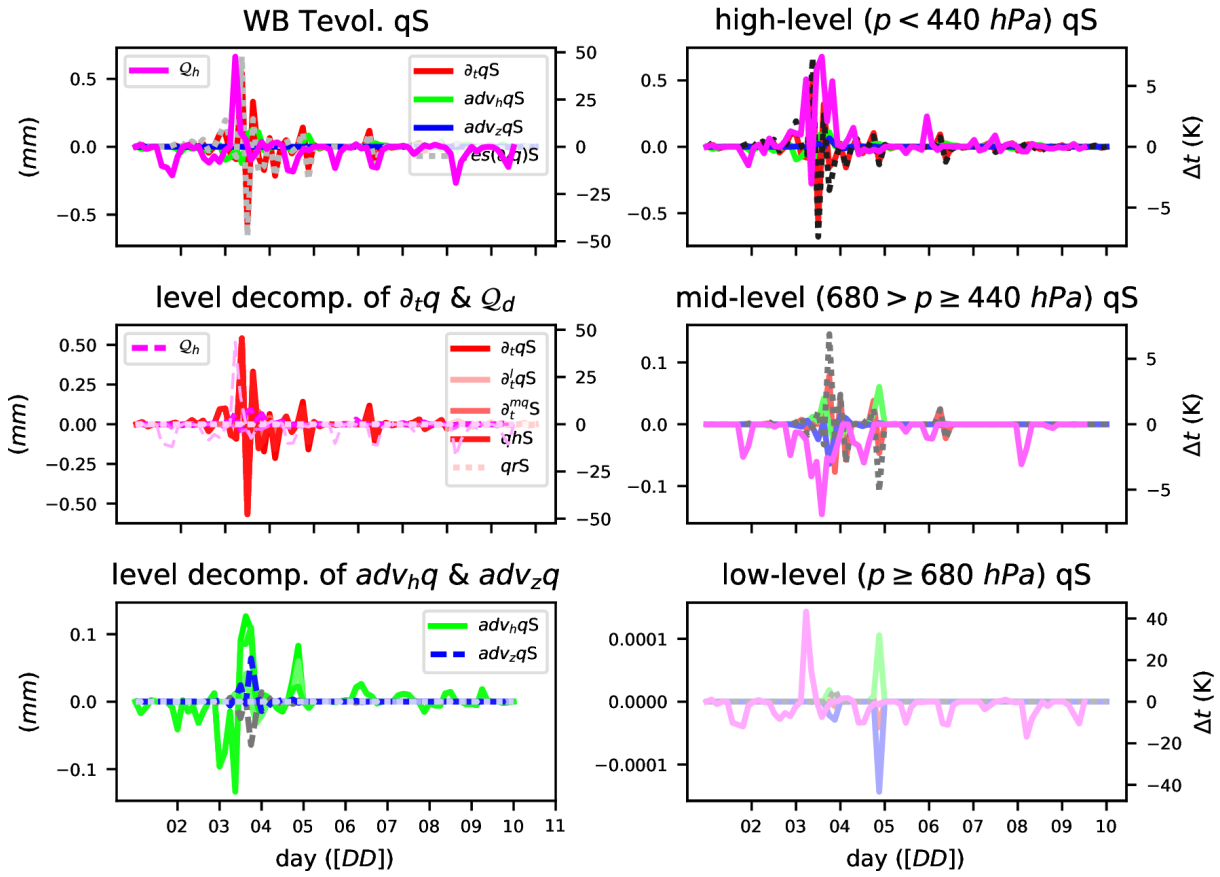


Figure 14. The same as in figure 12, but for snow

- **RUC** [`visibility_diag = 2`]: Visibility is computed using relative humidity (hur) as implemented in the RUC model (see equation 32 from Smirnova et al., 2000)

$$fogvis = 60.0 \exp \left[\frac{-2.5(hur \times 100 - 15)}{80} \right] \quad (32)$$

where hur : relative humidity (1, previously computed) and can be from the 2-m diagnostics or the first model layer. Visibility values are in km

5

- **FRAM-L** [`visibility_diag = 3`], (default): Visibility is computed using relative humidity (hur) after (see equation 33 from Gultepe and Milbrandt, 2010). In this work, a probabilistic approach is proposed to compute the visibility in three different bins: 95% , 50% and 5% of probability to get certain visibility (for $rh > 30\%$). As a matter of compromise in the module, the calculation with the 50% of probability has been chosen as the preferred one. Therefore,

Water Budget terms on 2012-12-09 15 UTC

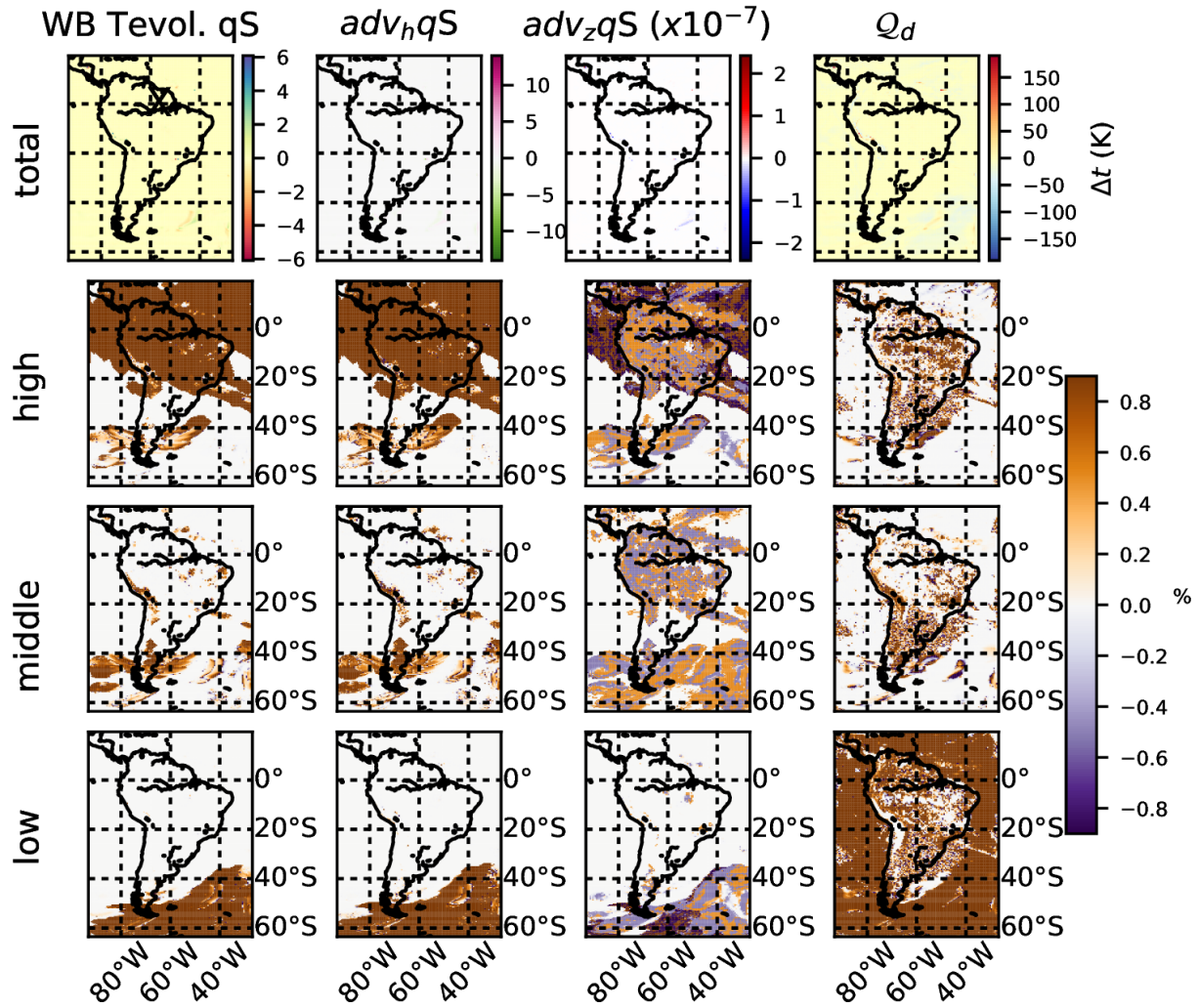


Figure 15. The same as in 13, bur for snow

this method provides the visibility that may occur with a 50% of probability.

$$f_{ogvis}^{prob} = \begin{cases} 95\% & -9.68 \times 10^{-14}hur^{7.19} + 52.20 \\ 50\% & -5.19 \times 10^{-10}hur^{5.44} + 40.10 \\ 5\% & -0.000114rh^{2.70} + 27.45 \end{cases} \quad (33)$$

where hur : relative humidity (1) and can be from 2-m diagnostics or first model layer. Visibility values are in km

5 Provided values of visibility during a fog event are: the minimum, maximum and mean values within output time steps (9freq) when fog occurred. Different choices are controlled throughout namelist variables: `visibility_diag` determines the method used to compute visibility, `fogvars` determines the source of the relative humidity to be used as input in the visibility method. User can choose between the relative humidity from the first model layer (hur) `fogvars=1` (default value) or from the 2-m diagnostics ($hurs$) `fogvars=2`. Some preliminary results of an extreme fog episode in central Argentina are provided in figure 16. Results strongly differ among fog implementations. The best agreement with a satellite visible channel picture for a given time of the event is obtained when the default setting is used ('FRAM-L' method with 'hur' values as input).

10 It is known that certain methods for calculating visibility rely on numerical adjustments on certain observational data taken under certain circumstances and at specific places (e.g.: for FRAM-L adjusted values come from observations from a Canadian airport). It would be desirable to provide a more generic "all places/purposes" approach (if possible). It is recommended to take this variable with a certain care.

15 **tfog: time of presence of fog**

Fog can be diagnosed when the visibility is lower than 1 km (WMO, 2010b). `tfog` accounts for the period during which the grid point has visibility lower than 1 km during 9freq (see equation 34)

$$tfog = \sum_{it}^{9freq} \delta t, [vis(it) \leq 1km] \quad (34)$$

where vis : visibility (km) below 1 km. δt : model time step (s)

20 **5 Optimization**

Regional climate dynamical downscaling experiments like the ones under the scope of CORDEX require long continuous simulations which consume larger amounts of HPC resource for a long period of time. Therefore, a series of tests were carried out in order to investigate the impact on the time of integration when module is activated. First version of the module (v1.0) was known to introduce about 40% decrease in time step speed of integration (highly dependent on HPC, model configuration and domain specifications). In order to improve model performance when the module is activated, the module was upgraded to the version v1.1. Since this new version, a series of optimizations of the code and pre-compilation flags were activated (CDXWRF). Following this implementation (instead via regular WRF namelist options), two main goals were achieved: (1) the

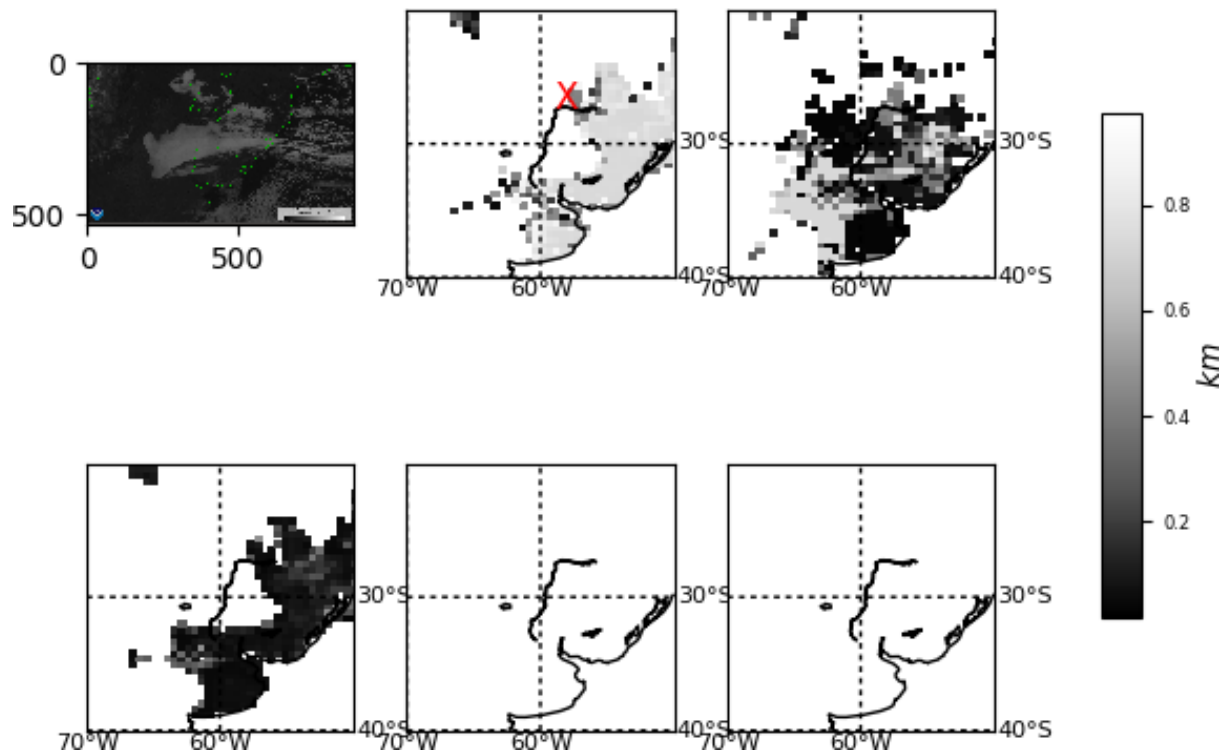
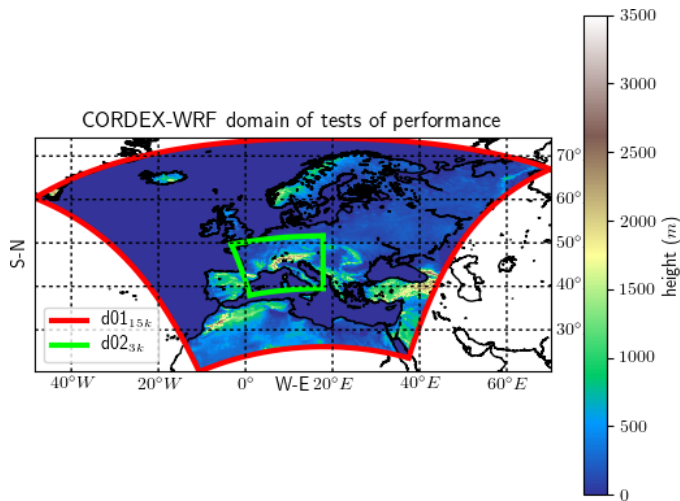


Figure 16. On 2007 June 30th at 12, comparison of the different configurations of the diagnostics of the mean fog visibility (in 1 hour) to the satellite image from GOES-12 (top left) at the same time in the visible channel (courtesy of NOAA-CLASS), default vis3vars1 (fogvisibility=3, fogvars=1; top middle), vis3vars2 (top right), vis1vars1 (bottom left), vis2vars1 (bottom middle) and vis2vars2 (bottom right)

amount of variables kept in memory during model execution was reduced and (2) the number of conditions (mainly avoiding IF statements) to be checked and calculations at each execution step of integration were reduced as well.

A domain as shown in figure 17 has been set-up to perform short runs (5 days) to check the changes in performance of the WRF model when the module is activated in its different possible configurations. In order to avoid non-homogeneous communication among the cluster nodes (which would affect the analysis), all the simulations were executed on a single node and with the WRF model compiled with intel and GNU Fortran compilers. Tests are performed at the HPC ‘Fram’ (https://www.sigma2.no/Fram) from the Norwegian academic HPC infrastructure. Fram is based on Lenovo NeXtScale nx360, constituted by CPU types: Intel E5-2683v4 2.1 GHz and Intel E7-4850v4 2.1 GHz.

The execution time is calculated as the mean elapsed time used during the 5-day model integration. Elapsed time necessary for each simulation step is available from the standard output of the model run (rs1.error.0000 WRF ASCII file). In this file, WRF users can get the elapsed time for all the time steps of the model and domains of simulation. Different peaks of slower time steps (see in figure 18) coincide with Input/Output file operations, difference between day and night regimes and



- projection: rotated lat-lon
- $ref_{lat,lon}$: $N 38.05220^\circ, 1.07623^\circ E$
- $ref_{x,y}$: 118, 99,
- $pole_{lat,lon}$: $N 39.25000^\circ, 18.00000^\circ E$
- standard longitude: $18.00000^\circ W$
- grid-points (d01, d02): $360 \times 351, 491 \times 441$
- resolution (d01, d02): $0.1375 \times 0.1375^\circ, 0.0275 \times 0.0275^\circ$
- time step (d01, d02): 75, 15 s
- d01 grid point ref. for d02: 118, 99
- vertical levels: 50

Figure 17. 2-nested domain WRF3.8.1 configuration where different performance tests were carried out.

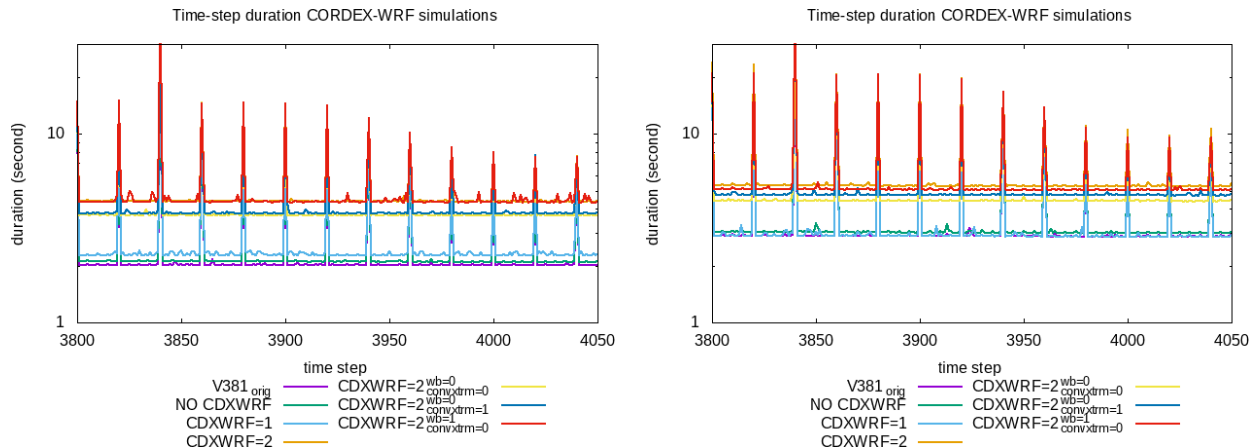


Figure 18. Elapsed times for each individual time step integration on nested domain d02 (time steps from number 3800 [simulating date 2014-01-01 15:35:00 UTC] to 4050 [2014-01-01 16:37:30 UTC]). Model was ran with different module configurations. See text for more details. Larger time steps are related to activation of the short/long-wave radiation scheme (every 5 minutes). For WRF compilation using ‘ifort’ (left) and ‘gcc’ compilers (right)

when different physical schemes (mainly the radiative scheme) are activated on a given frequency (e.g. `radt`). For a simulation covering 5-days with a time steps of 15 seconds, one obtains 28800 time steps. The sample of 28800 time steps is considered to be large enough to be representative for the mean time-step of the whole simulation.

Table 4 describes the different configurations and namelist options used in this performance test. The first simulation (labeled 5 V381orig), which is used as reference, is the simulation with the original version of the model (here version 3.8.1) without

Table 4. Mean elapsed time step of 5-day simulation and difference in time with respect original version of the code (v381orig) for different model configurations. See text for more details

label	description	ifort		gcc	
		$\langle t_{step} \rangle$ (s)	diff. (%)	$\langle t_{step} \rangle$ (s)	diff. (%)
v381orig	original WRF 3.8.1	2.4248	-	3.5174	-
NOCDXWRF	without CDXWRF	2.5058	3.34	3.6486	3.73
CDXWRF1	CDXWRF=1	2.6938	11.09	3.5070	-0.27
CDXWRF2	CDXWRF=2	4.8296	99.17	5.9958	70.46
CDXWRF2_00	CDXWRF=2 wb_diag=0 & convxtrm_diag=0	4.2038	73.37	5.0736	44.24
CDXWRF2_01	CDXWRF=2 wb_diag=0 & convxtrm_diag=1	4.2388	74.81	5.4120	53.86
CDXWRF2_10	CDXWRF=2 wb_diag=1 & convxtrm_diag=0	4.8510	100.06	5.7534	63.57

the module. The other simulations are: activation of the module (setting CORDEXDIAG) without setting of the pre-compilation parameter CDXWRF (labeled NOCDXWRF); with pre-compilation parameter CDXWRF=1 (CDXWRF1); with pre-compilation parameter CDXWRF=2 and computing all extra calculations (CDXWRF2). Finally three more simulations (with CDXWRF=2) are made: without any extra calculation (CDXWRF2_00), without calculation of extra water-budget terms (CDXWRF2_01) and without extremes from convection indices (CDXWRF2_10). Results might present some inconsistencies due to the fact that certain computation of diagnostics depend on the stability at each grid point which might vary from run to run and work load of the HPC.

Results show that all the simulations (except CDXWRF1 with gcc) where the module has been activated are slower than the simulation with the original version of the code (v381orig, $\langle t \rangle = 2.4248$ [ifort], 3.5174[gcc] s). Simulation with version 1.3 of the module without pre-compilation flag CDXWRF (NOCDXWRF, $\langle t \rangle = 2.5058, 3.6486$ s) is the second fastest. Simulation becomes slower when all the extra calculations are performed (CDXWRF2 $\langle t \rangle = 4.8296, 5.9958$ s). The heaviest part of the module is related to the water budget computation (wb_diag=1), since when comparing to the simulation without extra calculations (CDXWRF2_00, $\langle t \rangle = 4.2038, 5.0736$ s) there is an increase of only about 1, 9% ($\langle t_{step} \rangle_{CDXWRF2_01}^{ifort} / \langle t_{step} \rangle_{CDXWRF2_00}^{ifort}, \langle t_{step} \rangle_{CDXWRF2_01}^{gcc} / \langle t_{step} \rangle_{CDXWRF2_00}^{gcc}$) of mean time step when only statistics of extreme convective indices is activated (CDXWRF2_01, $\langle t \rangle = 4.2388, 5.4120$ s), and 27, 19% ($\langle t_{step} \rangle_{CDXWRF2_10} / \langle t_{step} \rangle_{CDXWRF2_00} |_{ifort,gcc}$) when only water-budget terms are included (CDXWRF2_10, $\langle t \rangle = 4.8510, 5.7534$ s). Reduction on time-step for CDXWRF1 with gcc must be related to a moment where HPC ‘Fram’ experiences a period of very low working load.

These results are not conclusive (optimization of the module should be tested in other HPC resources, domains and compilers), but they provide a first insight on how the number of variables included during the integration (in the derived type grid) has an important effect on model performance by reducing/increasing the required amount of memory. CDXWRF1 and

CDXWRF2_00 perform the same amount of computations of diagnostics, but the mean time step in CDXWRF2_00 is almost doubled (because in the CDXWRF2_00 case all the additional variables are defined but not diagnosed).

6 Summary and outlook

A series of modifications have been introduced into the WRF model in order to make it more suitable for CORDEX experiments. These modifications include the creation of a new module for the model and the modification of certain sections of the code. With the module implementation, the users have the option to retrieve directly from the WRF output all the ‘Core’ and almost all the ‘Tier1’ CORDEX variables. Moreover, a series of additional variables not required by CORDEX which may be of a high interest to the regional climate modelling community have been included. Additionally a series of ‘generic’ variables are provided which are independent of the model set-up, which should allow more robust intercomparisons.

A WRF user participating in a CORDEX regional climate experiment will strongly benefit when activating the module presented here, since it avoids most of the complex and time consuming post-processing efforts necessary to obtain diagnostics required by CORDEX. Furthermore, the module makes use of the WRF online interpolation to pressure levels of certain 3-dimensional variables which is an expensive computational task. In addition, statistical values (like minimum, fluxes or accumulation) are computed using the actual values obtained during the model integration. Since the use of the module avoids the post-processing of the model output, there is no need to keep large amount of additional fields which extremely reduces the requirements of disk storage (roughly 50 % less of disk storage).

In order to reduce the impact of the module during model integration, set-up of the module is managed before the compilation of the model via pre-compilation flags. This reduces the requirements of memory since the calculation of only required variables is activated. Different tests showed that the model performance is significantly improved when the code and the variables are constrained to the required ones (and not all available included), and managed via pre-compilation flags (and via conditional statements from the model namelist as it is usually done in the WRF model).

The module also establishes a first attempt to provide and implement generic definitions of certain diagnostics directly from regular model fields. This specific effort is intended to provide diagnostics which would not depend on the model configuration or/and the model itself, which will make intercomparisons between models more accurate, generic and trustworthy. If there is also a certain coordination (e.g. within the CORDEX community) of the definition of these diagnostics, climate studies will benefit. Common definitions of diagnostics in a coordinated way across different modeling platforms, will ensure robustness of model intercomparisons.

There are certain variables from CORDEX ‘Tier1’ protocol, which are not yet introduced but required. These are: (1) snow melt, (2) the daily maximum wind speed of gust at 100 m (wsgsmax100) which is not yet introduced because of lack of an appropriate method, and (3) intra-cloud, ground and total lightning flashes (ic_lightning, cg_lightning, tot_lightning) are not yet introduced even though a lightning scheme is already implemented in WRF. However this implementation is not appropriate for small domain patches used during parallelization. It has been found (L. Fita and M. Álvarez, CIMA personal communication) that lightning flashes rates become patchy when size of the parallel domain partitions are smaller than the actual extent of the

cell. Lightning methods (Price and Rind, 1992; Wong et al., 2013) require to encompass the whole convective cell to analyze the updraft within the cell. When patches do not cover the whole cell, different values from the same cell at each parallel fraction of the domain is provided from which spatial inconsistencies arise.

The module provides almost all the required all the CORDEX variables. However, user still needs to perform some post-
5 processing of the output data in order to meet CORDEX standards. Mainly:

- Computation of the required different statistical values as daily, monthly and seasonal extremes (minimum, maximum, accumulations, means)
- Cmorization of the output understood as: 1 file per variable, right metadata and attributes and general CF-compliant standard specifications

10 There are some additional features that would make the module more useful for the climate regional community, which are also required by the CORDEX specifications. However, in order to prioritize the production of a basic, full working, and useful version of the module as soon as possible, following aspects are planned for the future iterations and updates:

1. Flexibility-enhancement: Capacity to provide a more flexible module easy to be maintained and updated with new requirements from CORDEX or WRF model updates
- 15 2. Introduce new variables with a special focus on the implementation of more ‘generic’ variable definitions
3. CF-compliant/cmorization output: WRF output does not fully follows CF-conventions. Thus a huge coding effort needs to be done in order to provide a full CF-compliant output directly from it. User still needs to process the output of the model in order to provide data following all the CORDEX guidelines. Due to uncovered steps of the CF-standard, a user of the WRF model still needs to: concatenate files, change names and attributes of variables, calculate temporal
20 statistics over different periods (daily, monthly, seasonal) and provide the right time-variables in order to fully reach the CF-standard which followed by CORDEX. However, these steps are computationally lighter and easier to perform in comparison to the computation of the different diagnostics and the vertical pressure interpolation already introduced in the module.

The incompatibility between WRF output and CF-convention can be overcome with the use of a complementary dedicated
25 I/O library. This has been done for example in the RegIPSL platform (<https://sourcesup.renater.fr/wiki/morcemed/Home>, which uses WRF as atmospheric model) which uses XIOS (<http://forge.ipsl.jussieu.fr/iomserver>) libraries to manage the I/O.

The module is currently being implemented in the repository of the code in a dedicated branch. Once the module is fully implemented in the latest version of the model, and some additional tests are made, it might be possible that model developing team decides to include it in the main source of the code.

30 *Code availability.* The discussed version of the module (1.3) in the present article for 4 different WRF versions (3.7.1 Fita, 2018a), (3.8.1 Fita, 2018b), (3.9.1.1 Fita, 2018c), (4.0 Fita, 2018d) is available throughout Digital Object Identifiers (DOI):

- WRF-CORDEX module version 1.3 for WRFV3.7.1: DOI 10.5281/zenodo.1469639 (e.g. <https://doi.org/10.5281/zenodo.1469639>)
- WRF-CORDEX module version 1.3 for WRFV3.8.1: DOI 10.5281/zenodo.1469645
- WRF-CORDEX module version 1.3 for WRFV3.9.1.1: DOI 10.5281/zenodo.1469647
- WRF-CORDEX module version 1.3 for WRFV4.0: DOI 10.5281/zenodo.1469651

5 It is strongly recommended to make contact with lluis.fita@cima.fcen.uba.ar. This would facilitate the creation of a community of users of the module, and make it more easy to share the updates and improvements with the users.

A WIKI page has been set-up in order to provide detailed instructions how to manage the module, and to give information about upcoming module versions and improvements. The page will be regularly updated and can be found at:

<http://wiki.cima.fcen.uba.ar/mediawiki/index.php/CDXWRF>

Appendix B: Requested CORDEX variables

Here is provided a generic list of requested variables by CORDEX. Reader is advised that there is not a single CORDEX requirement variables list. It might depend on the experiment. However, hoping to provide a generic list of variables, a table with the CORDEX requirements in tables to is provided. The source of the table is from the ESGF servers at <https://www.earthsystemcog.org/doc/detail/1065/>. Same variable might appear at different levels (Core, Tier-1, Tier-2) as function of the requested frequency and/or if should be provided as statistical value between output frequency or instantaneous value, as well as, depending on the experiment (FPS_Alps experiment requested additional variables provided in table B4).

Appendix C: CORDEX variables in WRF

C1 Core variables

10 The variables given in Tables C1 and C2 are always provided when the module is activated with the use of the precompilation flag CORDEXDIAG. These variables appear in the auxiliary output file number 9. It is recommended to set the name of this file as `wrfcdx_d<domain>_<date>`. Note that some variables might not be produced according to chosen namelist options.

C2 Tier 1

15 The variables given in Table C3 are provided if the preprocessing variable CDXWRF is set to 1. These variables also appear in the auxiliary output file number 9. Note that some variables might not be produced according to chosen namelist options.

C3 Additional

The variables from Tables from C4 to C7 are provided if the preprocessing variable CDXWRF is set to 2. These variables also appear in the auxiliary output file number 9. Note that some variables might not be produced according to chosen namelist options.

20 C4 Instantaneous

The variables given in Table C8 are provided if some modifications are done in the WRF code. These variables represent instantaneous values for certain number of variables which are internally used. These variables also appear in the auxiliary output file number 9. Note that some variables might not be produced according to chosen namelist options.

Table B1. CORDEX requested variables (from <https://www.earthsystemcog.org/doc/detail/1065/>). *ovarn*: variable output name, *L*: level of requirement (c, core; 1, tier-1; 2, tier-2), *activation*: option to get the variable from WRF (wrfout, standard WRF output; clwrf, after clWRF modifications; cdxdia, basic CORDEXDIAG compilation, cdxwrf1, CORDEXDIAG activating pre-compilation flag cdxwrf=1; cdxwrf2, CORDEXDIAG activating pre-compilation flag cdxwrf=2; not available, not retrievable in this version of the module), ^a indicates when a variable has a dependency on the chosen scheme

#	<i>ovarn</i>	<i>units</i>	<i>L</i>	<i>Variable comment</i>	<i>activation</i>
1	tas	K	c	2-metre Air Temperature	wrfout
2	tasmax	K	c	Daily-Maximum 2-metre Air Temperature	clwrf
3	tasmin	K	c	Daily-Minimum 2-metre Air Temperature	clwrf
4	pr	kg m-2 s-1	c	Precipitation	cdxdia
5	ps	Pa	1	Surface Pressure	cdxdia
6	psl	Pa	c	Mean Sea Level Pressure	cdxdia
7	huss	1	c	2-metre Specific Humidity	cdxdia
8	sfcWind	m s-1	c	10-metre Wind Speed	cdxdia
9	sfcWindmax	m s-1	c	Daily-Maximum 10-metre Wind Speed	clwrf
10	clt ^a	%	c	Total Cloud Cover	cdxdia
11	sund	s	c	Sunshine Hours	cdxdia
12	rsds ^a	W m-2	c	Surface Downwelling Shortwave Radiation	cdxdia
13	rlds ^a	W m-2	c	Surface Downwelling Longwave Radiation	cdxdia
14	hfls ^a	W m-2	c	Surface Latent Heat Flux	cdxdia
15	hfss ^a	W m-2	c	Surface Sensible Heat Flux	cdxdia
16	rsus ^a	W m-2	c	Surface Upwelling Shortwave Radiation	cdxdia
17	rlus ^a	W m-2	c	Surface Upwelling Longwave Radiation	cdxdia
18	evspsbl ^a	kg m-2 s-1	c	Surface Evaporation	cdxdia
19	evspsblpot	kg m-2 s-1	1	Potential Evapotranspiration	cdxdia
20	mrfso	kg m-2	c	Soil Frozen Water Content	not available
21	mrros	kg m-2 s-1	c	Surface Runoff	cdxdia
22	mrro	kg m-2 s-1	c	Total Runoff	cdxdia
23	mrso	kg m-2	c	Total Soil Moisture Content	cdxdia

^ascheme dependant

Table B2. Continuation of table B1

24	snw	kg m-2	c	Snow Amount	cdxdiag
25	snm	kg m-2 s-1	c	Snow Melt	not available
26	prhmax	kg m-2 s-1	1	Maximum 1-hour Precipitation Rate within 24 hour period	not available
27	prc	kg m-2 s-1	1	Convective Precipitation	cdxdiag
28	rlut ^a	W m-2	c	TOA Outgoing Longwave Radiation	cdxdiag
29	rsdt ^a	W m-2	c	TOA Incident Shortwave Radiation	cdxdiag
30	rsut ^a	W m-2	c	TOA Outgoing Shortwave Radiation	cdxdiag
31	uas	m s-1	c	Eastward 10-metre Wind Speed	cdxdiag
32	vas	m s-1	c	Northward 10-metre Wind Speed	cdxdiag
33	wsgsmax	m s-1	1	Maximum 10-metre Gust Wind Speed	cdxdiag
34	tauu	Pa	1	Surface Downward Eastward Wind Stress	cdxdiag
35	tauv	Pa	1	Surface Downward Northward Wind Stress	cdxdiag
36	ts	K	1	Surface (Skin) Temperature	cdxdiag
37	zmla ^a	m	1	Atmospheric Boundary Layer Thickness	wrfout
38	prw	kg m-2	1	Column Water Vapour	cdxdiag
39	clwvi	kg m-2	1	Column Condensed (liquid+ice) Water Content	cdxdiag
40	clivi	kg m-2	1	Column Ice Water Content	cdxdiag
41	ua850	m s-1	c	Zonal (eastward) Wind at 850 hPa	cdxdiag
42	va850	m s-1	c	Meridional (northward) Wind at 850 hPa	cdxdiag
43	ta850	K	c	Temperature at 850 hPa	cdxdiag
44	hus850	1	c	Specific Humidity at 850 hPa	cdxdiag
45	ua500	m s-1	c	Zonal (eastward) Wind at 500 hPa	cdxdiag
46	va500	m s-1	c	Meridional (northward) Wind at 500 hPa	cdxdiag
47	zg500	m	c	Geopotential Height at 500 hPa	cdxdiag
48	ta500	K	c	Temperature at 500 hPa	cdxdiag
49	ua200	m s-1	c	Zonal (eastward) Wind at 200 hPa	cdxdiag
50	va200	m s-1	c	Meridional (northward) Wind at 200 hPa	cdxdiag
51	ta200	K	c	Temperature at 200 hPa	cdxdiag
52	zg200	m	c	Geopotential Height at 200 hPa	cdxdiag
53	clh ^a	%	1	High Clouds (p<440hPa)	cdxdiag
54	clm ^a	%	1	Medium Clouds (680hPa)	cdxdiag
55	cll ^a	%	1	Low Clouds (p>680hPa)	cdxdiag

^ascheme dependant

Table B3. Continuation of table B1

56	snc	%	c	Snow Area Fraction	cdxdiag
57	snd	m	c	Snow Depth	cdxdiag
58	sic	%	c	Sea Ice Fraction	not available
59	prsn	kg m-2 s-1	l	Snowfall Flux	cdxdiag
60	areacella	m2	c	Atmosphere Grid-Cell Area	not available
61	orog	m	c	Surface Altitude	not available
62	sftlf	%	c	Land Area Fraction	not available
63	sftgif	%	c	Fraction of Grid Cell Covered with Glacier	not available
64	mrsofc	kg m-2	c	Capacity of Soil to Store Water	not available
65	rootd	m	c	Maximum Root Depth	not available

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Table B4. As in B1, but the additional variables requested by the CORDEX FPS_Alps experiment

#	ovarn	units	L	Variable comment	CDXWRF option
1	ua1000	m s-1	c	Zonal (eastward) Wind at 1000 hPa	cdxdiag
2	va1000	m s-1	c	Meridional (northward) Wind at 1000 hPa	cdxdiag
3	ta1000	K	c	Temperature at 1000 hPa	cdxdiag
4	hus1000	1	c	Specific Humidity at 1000 hPa	cdxdiag
5	zg1000	m	c	Geopotential Height at 1000 hPa	cdxdiag
6	wa1000	m s-1	c	Vertical (upward) Wind at 1000 hPa	cdxdiag
7	ua925	m s-1	c	Zonal (eastward) Wind at 925 hPa	cdxdiag
8	va925	m s-1	c	Meridional (northward) Wind at 925 hPa	cdxdiag
9	ta925	K	c	Temperature at 925 hPa	cdxdiag
10	hus925	1	c	Specific Humidity at 925 hPa	cdxdiag
11	zg925	m	c	Geopotential Height at 925 hPa	cdxdiag
12	wa925	m s-1	c	Vertical (upward) Wind at 925 hPa	cdxdiag
13	ua700	m s-1	c	Zonal (eastward) Wind at 700 hPa	cdxdiag
14	va700	m s-1	c	Meridional (northward) Wind at 700 hPa	cdxdiag
15	ta700	K	c	Temperature at 700 hPa	cdxdiag
16	hus700	1	c	Specific Humidity at 700 hPa	cdxdiag
17	zg700	m	c	Geopotential Height at 700 hPa	cdxdiag
18	wa700	m s-1	c	Vertical (upward) Wind at 700 hPa	cdxdiag
19	wa850	m s-1	c	Vertical (upward) Wind at 850 hPa	cdxdiag
20	wa500	m s-1	c	Vertical (upward) Wind at 500 hPa	cdxdiag
21	wa200	m s-1	c	Vertical (upward) Wind at 200 hPa	cdxdiag
22	wsgsmax100m	m s-1	1	Maximum wind speed of gust at 100m	cdxdiag
23	ic_lightning	counts	1	Intra-cloud lightning flashes	not available
24	cg_lightning	counts	1	Cloud-ground lightning flashes	not available
25	total_lightning	counts	1	Total number lightning flashes	not available
26	clgvi	kg m-2	1	Vertical integral of mass content of graupel	cdxwrf=1
27	clhvi	kg m-2	1	Vertical integral of mass content of hail	cdxwrf=1
28	CAPE	J Kg-1	1	2-D Maximum convective available potential energy	cdxwrf=1
29	CIN	J Kg-1	1	2-D Maximum convective inhibition	cdxwrf=1

Table C1. Description of CORDEX Core variables provided with the module. kind specifies when the variable is computed; a: computed all time steps, o: only at output time, t: according to a frequency in the namelist, s: statistic value from internal integration values and initialized after each output time step

CF name	WRF name	description	units	kind
2D				
lon	LON	LONGITUDE	degrees_east	o
lat	LAT	LATITUDE	degrees_north	o
cltmean	CLTMEAN	MEAN TOTAL CLOUDINESS IN CORDEX OUTPUT	%	t
cllmean	CLLMEAN	MEAN LOW-LEVEL CLOUDINESS (p >= 68000 Pa) IN CORDEX OUTPUT	%	t
clmmean	CLMMEAN	MEAN MID-LEVEL CLOUDINESS (44000 <= p < 68000 Pa) IN CORDEX OUTPUT	%	t
clhmean	CLHMEAN	MEAN HIGH-LEVEL CLOUDINESS (p < 44000 Pa) IN CORDEX OUTPUT	%	t
mrso	MRSO	TOTAL SOIL CONTENT	kgm-2	o
prw	PRW	WATER VAPOR PATH	kgm-2	o
psl	PSL	SEA LEVEL PRESSURE	Pa	o
clwvi	CLWVI	LIQUID WATER PATH	kgm-2	o
clivi	CLIVI	ICE WATER PATH	kgm-2	o
hurs ^a	HURS	2M RELATIVE HUMIDITY	1	o
huss	HUSS	2M SPECIFIC HUMIDITY	1	o
slw	SLW	TOTAL SOIL LIQUID WATER CONTENT	kgm-2	o
uas	UAS	10M EASTWARD WIND SPEED	ms-1	o
vas	VAS	10M NORTHWARD WIND SPEED	ms-1	o
wsgsmax	WSGSMAX	Maximum near-surface wind speed of gust	ms-1	s
usgsmax	USGSMAX	Eastward maximum near-surface wind speed of gust	ms-1	s
vsgsmax	VSGSMAX	Northward maximum near-surface wind speed of gust	ms-1	s
totwsgsmax	TOTWSGSMAX	Total (TKE + h. pr) Maximum near-surface wind speed of gust	ms-1	s
totugsmax	TOTUGSMAX	Total Eastward maximum near-surface wind speed of gust	ms-1	s
totvgsmax	TOTVGSMAX	Total Northward maximum near-surface wind speed of gust	ms-1	s
wsz100max	WSZ100MAX	Maximum 100m nwind speed	ms-1	s
uz100max	UZ100MAX	Eastward maximum 100 m wind speed	ms-1	s
vz100max	VZ100MAX	Northward maximum 100 m wind speed	ms-1	s
sund	SUND	SUNSHINE LENGTH (ac. time SWDOWN > 120. Wm-2)	second	s

^aneeded for other variables

Table C2. Continuation of Table C1 of Core variables

tauu	TAUU	northward downward wind stress at 10 m	m2s-2	o
tauv	TAUV	eastward downward wind stress at 10 m	m2s-2	o
tauugen	TAUUGEN	generic eastward downward wind stress at 10 m	m2s-2	o
tauvgen	TAUVGEN	generic northward downward wind stress at 10 m	m2s-2	o
rsds	RSDS	mean surface Downwelling Shortwave Radiation	Wm-2	s
rlds	RLDS	mean surface Downwelling Longwave Radiation	Wm-2	s
hfls	HFLS	mean surface Upward Latent Heat Flux	Wm-2	s
hfss	HFSS	mean surface Upward Sensible Heat Flux	Wm-2	s
rsus	RSUS	mean surface Upwelling Shortwave Radiation	Wm-2	s
rlus	RLUS	mean surface Upwelling Longwave Radiation	Wm-2	s
rsusgen	RSUSGEN	mean generic surface Upwelling Shortwave Radiation	Wm-2	s
rlusgen	RLUSGEN	mean generic surface Upwelling Longwave Radiation	Wm-2	s
evspsbl	EVSPSBL	mean evaporation	kgm-2s-1	s
evspsblpot	EVSPSBLPOT	mean potential evapotranspiration	kgm-2s-1	s
evspsblpotgen	EVSPSBLPOTGEN	mean generic potential evapotranspiration	kgm-2s-1	s
cd	CDCDX	drag coefficient	-	o
cdgen	cdgen	generic drag coefficient	-	o
snc	SNC	mean snow area fraction	%	s
snd	SND	mean snow depth	m	s
mrros ^a	MRROS	mean surface Runoff	kgm-2s-1	s
mrro ^a	MRRO	mean total Runoff	kgm-2s-1	s
mrsol ^a	MRSOL	mean total water content of soil layer	kgm-2	s
pr	PR	precipitation flux	kgm-2s-1	s
prl	PRL	large scale precipitation flux	kgm-2s-1	s
prc	PRC	convective precipitation flux	kgm-2s-1	s
prsh	PRSH	shallow-cumulus precipitation flux	kgm-2s-1	s
prsn	PRSN	solid precipitation flux	kgm-2s-1	s
snw	SNW	accumulated snow precipitation	kgm-2	s
rsdt	RSDT	mean top of the atmosphere (TOA) incident shortwave radiation	kgm-2	s
rsut	RSUT	mean TOA outgoing shortwave radiation	kgm-2	s
rlut	RLUT	mean TOA outgoing Longwave radiation	kgm-2	s
ps	CDXPS	surface pressure	Pa	o
ts	CDXTS	skin temperature	K	o

^aunmasked to sea points

Table C3. As in C1, but for the description of CORDEX Tier1 variables provided with the module

CF name	WRF name	description	units	kind
clgvi	CLGVI	GRAUPEL WATER PATH	kgm-2	o
clhvi	CLHVI	HAIL WATER PATH	kgm-2	o
zmlagen	ZMLAGEN	Generic boundary layer height $\theta(zmlagen) > \min(\theta[mix. layer]) + 1.5K$	m	o
capemin ^a	CAPEMIN	MINIMUM CONVECTIVE AVAILABLE POTENTIAL ENERGY	Jkg-1	s
capemax ^a	CAPEMAX	MAXIMUM CONVECTIVE AVAILABLE POTENTIAL ENERGY	Jkg-1	s
capemean ^a	CAPEMEAN	MEAN CONVECTIVE AVAILABLE POTENTIAL ENERGY	Jkg-1	s
cinmin ^a	CINMIN	MINIMUM CONVECTIVE INHIBITION	Jkg-1	s
cinmax ^a	CINMAX	MAXIMUM CONVECTIVE INHIBITION	Jkg-1	s
cinmeann ^a	CINMEAN	MEAN CONVECTIVE INHIBITION	Jkg-1	s
lfcppmin ^a	LFCPPMIN	MINIMUM PRESSURE LEVEL FREE CONVECTION	Pa	s
lfcppmax ^a	LFCPPMAX	MAXIMUM PRESSURE LEVEL FREE CONVECTION	Pa	s
lfcppmean ^a	LFCPPMEAN	MEAN PRESSURE LEVEL FREE CONVECTION	Pa	s
lfczmin ^a	LFCZMIN	MINIMUM HEIGHT LEVEL FREE CONVECTION	m	s
lfczmax ^a	LFCZMAX	MAXIMUM HEIGHT LEVEL FREE CONVECTION	m	s
lfczmean ^a	LFCZMEAN	MEAN HEIGHT LEVEL FREE CONVECTION	m	s
limin ^a	LIMIN	MINIMUM LIFTED INDEX	1	s
limax ^a	LIMAX	MAXIMUM LIFTED INDEX	1	s
limean ^a	LIMEAN	MEAN LIFTED INDEX	1	s

^ait will be computed if namelist parameter `convxtrm_diag` is set to 1

References

- Barella-Ortiz, A., Polcher, J., Tuzet, A., and Laval, K.: Potential evaporation estimation through an unstressed surface-energy balance and its sensitivity to climate change, *Hydrol. Earth Syst. Sci.*, 17, 4625–4639, <https://doi.org/10.5194/hess-17-4625-2013>, <https://www.hydrol-earth-syst-sci.net/17/4625/2013/>, 2013.
- 5 Benjamin, S. G. and Miller, P. A.: An Alternative Sea Level Pressure Reduction and a Statistical Comparison of Geostrophic Wind Estimates with Observed Surface Winds, *Mon. Weather Rev.*, 118, 2099–2116, [https://doi.org/10.1175/1520-0493\(1990\)118<2099:AASLPR>2.0.CO;2](https://doi.org/10.1175/1520-0493(1990)118<2099:AASLPR>2.0.CO;2), 1990.
- Bergot, T., Terradellas, E., Cuxart, J., Mira, A., Liechti, O., Mueller, M., and Nielsen, N. W.: Intercomparison of Single-Column Numerical Models for the Prediction of Radiation Fog, *Journal of Applied Meteorology and Climatology*, 46, 504–521, <https://doi.org/10.1175/JAM2475.1>, <https://DOI.org/10.1175/JAM2475.1>, 2007.
- 10

Table C4. As in C1, but for the description of additional variables provided with the module

CF name	WRF name	description	units	kind
tdsmin	TDSMIN	minimum surface dew point temperature	K	s
tdsmax	TDSMAX	maximum surface dew point temperature	K	s
tdsmean	TDSMEAN	mean surface dew point temperature	K	s
tfog	TFOG	time of presence of fog	seconds	s
fogvisbltymax	FOGVISBLTYMAX	maximum of visibility inside fog	km	s
fogvisbltymax	FOGVISBLTYMAX	maximum of visibility inside fog	km	s
fogvisbltymeans	FOGVISBLTYMEAN	mean of visibility inside fog	km	s
3D				
hur ^a	HUR	AIR RELATIVE HUMIDITY	1	a
hus	HUS	AIR SPECIFIC HUMIDITY	1	o
zg ^a	ZG	AIR GEOPOTENTIAL HEIGHT	m	a
press ^a	PRESS	AIR PRESSURE	Pa	a
ta ^a	TA	AIR TEMPERATURE	K	a
ua	UA	EARTH ROTATED AIR EASTWARD WIND SPEED	ms-1	o
va	VA	EARTH ROTATED AIR NORTHWARD WIND SPEED	ms-1	o
Water-Budget ^b				
Q_{hac}	WBACDIABH	Water Budget column integrated and time accumulation of diabatic heating from Micro-Physics	K	s
$\partial_t q_{vac}$	WBACPW	Water Budget column integrated and time accumulated for water vapor content	mm	s
$\partial_t q_{cac}$	WBACPWC	Water Budget col. int. & time accumulated for cloud content	mm	s
$\partial_t q_{rac}$	WBACPWR	Water Budget col. int. & time accumulated for rain content	mm	s
$\partial_t q_{sac}$	WBACPWS	Water Budget col. int. & time accumulated for snow content	mm	s
$\partial_t q_{iac}$	WBACPWI	Water Budget col. int. & time accumulated for ice content	mm	s
$\partial_t q_{hac}$	WBACPWH	Water Budget col. int. & time accumulated for hail content	mm	s
$\partial_t q_{gac}$	WBACPWG	Water Budget col. int. & time accumulated for graupel content	mm	s

^aneeded for other variables

^bVariables will be computed if namelist parameter output_wb is set to 1

Table C5. Continuation of Table C4 of additional variables

$adv_h qvac$	WBACF	W.B. c-int. acc. hor. convergence of water vapour (+, conv.; -, div.)	mm	s
$adv_h qcac$	WBACFC	W.B. c-int. acc. hor. convergence of cloud (+, conv.; -, div.)	mm	s
$adv_h qrac$	WBACFR	W.B. c-int. acc. hor. convergence of rain (+, conv.; -, div.)	mm	s
$adv_h qsac$	WBACFS	W.B. c-int. acc. hor. convergence of snow (+, conv.; -, div.)	mm	s
$adv_h qiac$	WBACFI	W.B. c-int. acc. hor. convergence of ice (+, conv.; -, div.)	mm	s
$adv_h qhac$	WBACFH	W.B. c-int. acc. hor. convergence of hail (+, conv.; -, div.)	mm	s
$adv_h qgac$	WBACFG	W.B. c-int. acc. hor. convergence of graupel (+, conv.; -, div.)	mm	s
$adv_z qvac$	WBACZ	W.B. c-int. acc. ver. convergence of water vapour (+, conv.; -, div.), always 0	mm	s
$adv_z qcac$	WBACZC	W.B. c-int. acc. ver. convergence of cloud (+, conv.; -, div.)	mm	s
$adv_z qrac$	WBACZR	W.B. c-int. acc. ver. convergence of rain (+, conv.; -, div.)	mm	s
$adv_z qsac$	WBACZS	W.B. c-int. acc. ver. convergence of snow (+, conv.; -, div.)	mm	s
$adv_z qiac$	WBACZI	W.B. c-int. acc. ver. convergence of ice (+, conv.; -, div.)	mm	s
$adv_z qhac$	WBACZH	W.B. c-int. acc. ver. convergence of hail (+, conv.; -, div.)	mm	s
$adv_z qgac$	WBACZG	W.B. c-int. acc. ver. convergence of graupel (+, conv.; -, div.)	mm	s
$Q_h^l ac$	WBACDIABHL	W.B. low level acc. of diabatic heating from MP	K	s
$Q_h^m ac$	WBACDIABHM	W.B. mid-level acc. of diabatic heating from MP	K	s
$Q_h^m ac$	WBACDIABHH	W.B. high-level acc. of diabatic heating from MP	K	s
$\partial_t^l qvac$	WBACPWL	W.B. low level ($p \geq 68000$ Pa) acc. for QV	mm	s
$\partial_t^m qvac$	WBACPWMV	W.B. mid level ($44000 \text{ Pa} \leq p < 68000 \text{ Pa}$) acc. for QV	mm	s
$\partial_t^h qvac$	WBACPWHV	W.B. high level ($p < 44000$ Pa) acc. for QV	mm	s
$adv_h^l qvac$	WBACFLV	W.B. low-lev. acc. hor. convergence of QV	mm	s
$adv_h^m qvac$	WBACFMV	W.B. mid-lev. acc. hor. convergence of QV	mm	s
$adv_h^h qvac$	WBACFHV	W.B. high-lev. acc. hor. convergence of QV	mm	s
$adv_z^l qvac$	WBACZLV	W.B. low level acc. ver. convergence of QV	mm	s
$adv_z^m qvac$	WBACZMV	W.B. mid level acc. ver. convergence of QV	mm	s
$adv_z^h qvac$	WBACZHV	W.B. high level acc. ver. convergence of QV	mm	s

Table C6. 2nd continuation of Table C4 of additional variables

$\partial_t^l \text{qcac}$	WBACPWLC	W.B. low level ($p \geq 68000$ Pa) acc. for QC	mm	s
$\partial_t^m \text{qcac}$	WBACPWMC	W.B. mid level ($44000 \text{ Pa} \leq p < 68000 \text{ Pa}$) acc. for QC	mm	s
$\partial_t^h \text{qcac}$	WBACPWHC	W.B. high level ($p < 44000$ Pa) acc. for QC	mm	s
$\text{adv}_h^l \text{qcac}$	WBACFLC	W.B. low-lev. acc. hor. convergence of QC	mm	s
$\text{adv}_h^m \text{qcac}$	WBACFMC	W.B. mid-lev. acc. hor. convergence of QC	mm	s
$\text{adv}_h^h \text{qcac}$	WBACFHC	W.B. high-lev. acc. hor. convergence of QC	mm	s
$\text{adv}_z^l \text{qcac}$	WBACZLC	W.B. low level acc. ver. convergence of QC	mm	s
$\text{adv}_z^m \text{qcac}$	WBACZMC	W.B. mid level acc. ver. convergence of QC	mm	s
$\text{adv}_z^h \text{qcac}$	WBACZHC	W.B. high level acc. ver. convergence of QC	mm	s
$\partial_t^l \text{qrac}$	WBACPWLR	W.B. low level ($p \geq 68000$ Pa) acc. for QR	mm	s
$\partial_t^m \text{qrac}$	WBACPWMR	W.B. mid level ($44000 \text{ Pa} \leq p < 68000 \text{ Pa}$) acc. for QR	mm	s
$\partial_t^h \text{qrac}$	WBACPWHR	W.B. high level ($p < 44000$ Pa) acc. for QR	mm	s
$\text{adv}_h^l \text{qrac}$	WBACFLR	W.B. low-lev. acc. hor. convergence of QR	mm	s
$\text{adv}_h^m \text{qrac}$	WBACFMR	W.B. mid-lev. acc. hor. convergence of QR	mm	s
$\text{adv}_h^h \text{qrac}$	WBACFHR	W.B. high-lev. acc. hor. convergence of QR	mm	s
$\text{adv}_z^l \text{qrac}$	WBACZLR	W.B. low level acc. ver. convergence of QR	mm	s
$\text{adv}_z^m \text{qrac}$	WBACZMR	W.B. mid level acc. ver. convergence of QR	mm	s
$\text{adv}_z^h \text{qrac}$	WBACZHR	W.B. high level acc. ver. convergence of QR	mm	s
$\partial_t^l \text{qsac}$	WBACPWLS	W.B. low level ($p \geq 68000$ Pa) acc. for QS	mm	s
$\partial_t^m \text{qsac}$	WBACPWMS	W.B. mid level ($44000 \text{ Pa} \leq p < 68000 \text{ Pa}$) acc. for QS	mm	s
$\partial_t^h \text{qsac}$	WBACPWHS	W.B. high level ($p < 44000$ Pa) acc. for QS	mm	s
$\text{adv}_h^l \text{qsac}$	WBACFLS	W.B. low-lev. acc. hor. convergence of QS	mm	s
$\text{adv}_h^m \text{qsac}$	WBACFMS	W.B. mid-lev. acc. hor. convergence of QS	mm	s
$\text{adv}_h^h \text{qsac}$	WBACFHS	W.B. high-lev. acc. hor. convergence of QS	mm	s
$\text{adv}_z^l \text{qsac}$	WBACZLS	W.B. low level acc. ver. convergence of QS	mm	s
$\text{adv}_z^m \text{qsac}$	WBACZMS	W.B. mid level acc. ver. convergence of QS	mm	s
$\text{adv}_z^h \text{qsac}$	WBACZHS	W.B. high level acc. ver. convergence of QS	mm	s
$\partial_t^l \text{qiac}$	WBACPWLI	W.B. low level ($p \geq 68000$ Pa) acc. for QI	mm	s
$\partial_t^m \text{qiac}$	WBACPWMI	W.B. mid level ($44000 \text{ Pa} \leq p < 68000 \text{ Pa}$) acc. for QI	mm	s
$\partial_t^h \text{qiac}$	WBACPWHI	W.B. high level ($p < 44000$ Pa) acc. for QI	mm	s
$\text{adv}_h^l \text{qiac}$	WBACFLI	W.B. low-lev. acc. hor. convergence of QI	mm	s
$\text{adv}_h^m \text{qiac}$	WBACFMI	W.B. mid-lev. acc. hor. convergence of QI	mm	s
$\text{adv}_h^h \text{qiac}$	WBACFHI	W.B. high-lev. acc. hor. convergence of QI	mm	s
$\text{adv}_z^l \text{qiac}$	WBACZLI	W.B. low level acc. ver. convergence of QI	mm	s
$\text{adv}_z^m \text{qiac}$	WBACZMI	W.B. mid level acc. ver. convergence of QI	mm	s
$\text{adv}_z^h \text{qiac}$	WBACZHI	W.B. high level acc. ver. convergence of QI	mm	s

Table C7. 2nd continuation of Table C4 of additional variables

$\partial_t^l \text{qgac}$	WBACPWLG	W.B. low level ($p \geq 68000$ Pa) acc. for QG	mm	s
$\partial_t^m \text{qgac}$	WBACPWMG	W.B. mid level ($44000 \text{ Pa} \leq p < 68000 \text{ Pa}$) acc. for QG	mm	s
$\partial_t^h \text{qgac}$	WBACPWHG	W.B. high level ($p < 44000$ Pa) acc. for QG	mm	s
$\text{adv}_h^l \text{qgac}$	WBACFLG	W.B. low-lev. acc. hor. convergence of QG	mm	s
$\text{adv}_h^m \text{qgac}$	WBACFMG	W.B. mid-lev. acc. hor. convergence of QG	mm	s
$\text{adv}_h^h \text{qgac}$	WBACFHG	W.B. high-lev. acc. hor. convergence of QG	mm	s
$\text{adv}_z^l \text{qgac}$	WBACZLG	W.B. low level acc. ver. convergence of QG	mm	s
$\text{adv}_z^m \text{qgac}$	WBACZMG	W.B. mid level acc. ver. convergence of QG	mm	s
$\text{adv}_z^h \text{qgac}$	WBACZHG	W.B. high level acc. ver. convergence of QG	mm	s
$\partial_t^l \text{qhac}$	WBACPWLH	W.B. low level ($p \geq 68000$ Pa) acc. for QH	mm	s
$\partial_t^m \text{qhac}$	WBACPWMH	W.B. mid level ($44000 \text{ Pa} \leq p < 68000 \text{ Pa}$) acc. for QH	mm	s
$\partial_t^h \text{qhac}$	WBACPWHH	W.B. high level ($p < 44000$ Pa) acc. for QH	mm	s
$\text{adv}_h^l \text{qhac}$	WBACFLH	W.B. low-lev. acc. hor. convergence of QH	mm	s
$\text{adv}_h^m \text{qhac}$	WBACFMH	W.B. mid-lev. acc. hor. convergence of QH	mm	s
$\text{adv}_h^h \text{qhac}$	WBACFHH	W.B. high-lev. acc. hor. convergence of QH	mm	s
$\text{adv}_z^l \text{qhac}$	WBACZLH	W.B. low level acc. ver. convergence of QH	mm	s
$\text{adv}_z^m \text{qhac}$	WBACZMH	W.B. mid level acc. ver. convergence of QH	mm	s
$\text{adv}_z^h \text{qhac}$	WBACZHH	W.B. high level acc. ver. convergence of QH	mm	s

Brasseur, O.: Development and Application of a Physical Approach to Estimating Wind Gusts, *Monthly Weather Review*, 129, 5–25, [https://doi.org/10.1175/1520-0493\(2001\)129<0005:DAAOAP>2.0.CO;2](https://doi.org/10.1175/1520-0493(2001)129<0005:DAAOAP>2.0.CO;2), [https://DOI.org/10.1175/1520-0493\(2001\)129<0005:DAAOAP>2.0.CO;2](https://DOI.org/10.1175/1520-0493(2001)129<0005:DAAOAP>2.0.CO;2), 2001.

5 Businger, J. A., Wyngaard, J. C., Izumi, Y., and Bradley, E. F.: Flux-Profile Relationships in the Atmospheric Surface Layer, *J. Atmos. Sci.*, 28, 181–189, [https://doi.org/10.1175/1520-0469\(1971\)028<0181:FPRITA>2.0.CO;2](https://doi.org/10.1175/1520-0469(1971)028<0181:FPRITA>2.0.CO;2), [https://doi.org/10.1175/1520-0469\(1971\)028<0181:FPRITA>2.0.CO;2](https://doi.org/10.1175/1520-0469(1971)028<0181:FPRITA>2.0.CO;2), 1971.

10 Coppola, E., Sobolowski, S., Pichelli, E., Raffaele, F., Ahrens, B., Anders, I., Ban, N., Bastin, S., Belda, M., Belusic, D., Caldas-Alvarez, A., Cardoso, R. M., Davolio, S., Dobler, A., Fernández, J., Fita, L., Fumiere, Q., Giorgi, F., Goergen, K., Guettler, I., Halenka, T., Heinzeller, D., Hodnebrog, O., Jacob, D., Kartsios, S., Katragkou, E., Kendon, E., Khodayar, S., Kunstmann, H., Knist, S., Lavín, A., Lind, P., Lorenz, T., Maraun, D., Marelle, L., van Meijgaard, E., Milovac, J., Myhre, G., H.-J.Panitz, Piazza, M., Raffa, M., Raub, T., Rockel, B., Schär, C., Sieck, K., Soares, P. M. M., Somot, S., Srnec, L., Stocchi, P., Tölle, M., Truhetz, H., Vautard, R., de Vries, H., and Warrach-Sagi, K.: A first-of-its-kind multi-model convection permitting ensemble for investigating convective phenomena over Europe and the Mediterranean, *Clim. Dyn.*, e-print first, <https://doi.org/10.1007/s00382-018-4521-8>, 2018.

15 de Rosnay, P., Polcher, J., Bruen, M., and Laval, K.: Impact of a physically based soil water flow and soil-plant interaction representation for modeling large-scale land surface processes, *J. Geophys. Res. (Atmospheres)*, 107(D11), <https://doi.org/10.1029/2001JD000634>, 2002.

Table C8. As in C1, but for the description of CORDEX ‘instantaneous’ variables provided with the module

CF name	WRF name	description	units	kind
clt	CLT	TOTAL CLOUDINESS	1	t
cll	CLL	LOW-LEVEL CLOUDINESS (p >= 68000 Pa)	1	t
clm	CLM	MID-LEVEL CLOUDINESS (44000 <= p < 68000 Pa)	1	t
clh	CLH	HIGH-LEVEL CLOUDINESS (p < 44000 Pa)	1	t
cape	CDXCAPE	CONVECTIVE AVAILABLE POTENTIAL ENERGY	Jkg-1	a/o ^a
cin	CIN	CONVECTIVE INHIBITION	Jkg-1	a/o ^a
lfcv	LFCV	PRESSURE LEVEL FREE CONVECTION	Pa	a/o ^a
lfcz	LFCZ	HEIGHT LEVEL FREE CONVECTION	m	a/o ^a
li	LI	LIFTED INDEX	1	a/o ^a
wsgs	WSGS	near-surface wind speed of gust	ms-1	a
usgs	USGS	Eastward near-surface wind speed of gust	ms-1	a
vsgs	VSGS	Northward near-surface wind speed of gust	ms-1	a
wsgspercen	WSGSPERCEN	Percentage of time steps where grid point got wind gust	%	s
totwsgspercen	TOTWSGSPERCEN	Percentage of time steps where grid point got total wind gust	%	s
wsz100	WSZ100	100m wind speed	ms-1	a
uz100	UZ100	Eastward 100 m wind speed	ms-1	a
vz100	VZ100	Northward 100 m wind speed	ms-1	a
fog	FOG	Whether there is fog (1: yes [vis < 1km]; 0: not)	-	a
fogvisblty	FOGVISBLTY	visibility inside fog	km	a
tds	TDS	surface dew point temperature	K	a

^adepending on namelist parameter convxtrm_diag 0:o, 1:a

Domínguez, M., Romera, R., Sánchez, E., Fita, L., Fernández, J., Jiménez-Guerrero, P., Montávez, J. P., Cabos, W. D., Liguori, G., and Gaertner, M. A.: Present climate precipitation and temperature extremes over Spain from a set of high resolution RCMs, *Climate research*, 58, 149–164, <https://doi.org/10.3354/cr01186>, 2013.

5 Evans, J. P., Ji, F., Lee, C., Smith, P., Argüeso, D., and Fita, L.: A regional climate modelling projection ensemble experiment - NARCLiM, *Geoscientific Model Development*, 7(2), 621–629, <https://doi.org/10.5194/gmd-7-621-2014>, 2014.

Fita, L.: WRF-CORDEX module version 1.3 for WRFV3.7.1, <https://doi.org/10.5281/zenodo.1469639>, <https://doi.org/10.5281/zenodo.1469639>, 2018a.

Fita, L.: WRF-CORDEX module version 1.3 for WRFV3.8.1, <https://doi.org/10.5281/zenodo.1469645>, <https://doi.org/10.5281/zenodo.1469645>, 2018b.

10 Fita, L.: WRF-CORDEX module version 1.3 for WRFV3.9.1.1, <https://doi.org/10.5281/zenodo.1469647>, <https://doi.org/10.5281/zenodo.1469647>, 2018c.

Table C9. Continuation of Table C8

3D-Water budget				
$qv_t tend$	QVTTEND	inter time step water vapor tendency	kgkg-1s-1	a
$qc_t tend$	QCTTEND	inter time step cloud tendency	kgkg-1s-1	a
$qr_t tend$	QRTTEND	inter time step rain tendency	kgkg-1s-1	a
$qs_t tend$	QSTTEND	inter time step snow tendency	kgkg-1s-1	a
$qi_t tend$	QITTEND	inter time step ice tendency	kgkg-1s-1	a
$qh_t tend$	QHTTEND	inter time step hail tendency	kgkg-1s-1	a
$qg_t tend$	QGTTEND	inter time step graupel tendency	kgkg-1s-1	a
$qv_h adv$	QV_HADV	Instantaneous QV Horizontal advection	kgkg-1	a
$qc_h adv$	QC_HADV	Instantaneous QC Horizontal advection	kgkg-1	a
$qr_h adv$	QR_HADV	Instantaneous QR Horizontal advection	kgkg-1	a
$qs_h adv$	QS_HADV	Instantaneous QS Horizontal advection	kgkg-1	a
$qi_h adv$	QI_HADV	Instantaneous QI Horizontal advection	kgkg-1	a
$qh_h adv$	QH_HADV	Instantaneous QH Horizontal advection	kgkg-1	a
$qg_h adv$	QG_HADV	Instantaneous QG Horizontal advection	kgkg-1	a
$qv_z adv$	QV_ZADV	Instantaneous QV Vertical advection	kgkg-1	a
$qc_z adv$	QC_ZADV	Instantaneous QC Vertical advection	kgkg-1	a
$qr_z adv$	QR_ZADV	Instantaneous QR Vertical advection	kgkg-1	a
$qs_z adv$	QS_ZADV	Instantaneous QS Vertical advection	kgkg-1	a
$qi_z adv$	QI_ZADV	Instantaneous QI Vertical advection	kgkg-1	a
$qh_z adv$	QH_ZADV	Instantaneous QH Vertical advection	kgkg-1	a
$qg_z adv$	QG_ZADV	Instantaneous QG Vertical advection	kgkg-1	a

Fita, L.: WRF-CORDEX module version 1.3 for WRFV4.0, <https://doi.org/10.5281/zenodo.1469651>, <https://doi.org/10.5281/zenodo.1469651>, 2018d.

5 Fita, L. and Flaounas, E.: Medicanes as subtropical cyclones: the December 2005 case from the perspective of surface pressure tendency diagnostics and atmospheric water budget, *Q. J. Royal Meteor. Soc.*, 144, 1028–1044, <https://doi.org/10.1002/qj.3273>, qJ-17-0198.R2, 2018.

Fita, L., Fernández, J., and García-Díez, M.: CLWRF: WRF modifications for regional climate simulation under future scenarios, *Proceedings of 11th WRF Users' Workshop*, 2010.

Fu, C., Wang, S., Xiong, Z., Gutowski, W. J., Lee, D. K., McGregor, J. L., Sato, Y., Kato, H., Kim, J. W., and Suh, M. S.: Regional climate model intercomparison project for Asia, *Bull. Amer. Meteor. Soc.*, 86(2), 257–266, 2005.

10 García-Díez, M., Fernández, J., Fita, L., and Yagüe, C.: Seasonal dependence of WRF model biases and sensitivity to PBL schemes over Europe, *Q. J. of Roy. Met. Soc.*, 139, 501–514, <https://doi.org/10.1002/qj.1976>, 2013.

Garratt, J.: *The Atmospheric Boundary Layer*, Cambridge Univ. Press, Cambridge, U.K., 1992.

- Geleyn, J. F. and Hollingsworth, A.: An economical analytical method for the computation of the interaction between scattering and line absorption of radiation, *Contrib. Atmos. Phys*, 52, 1–16, 1979.
- Giorgi, F. and Gutowski, W. J.: Regional Dynamical Downscaling and the CORDEX Initiative, *Annual Review of Environment and Resources*, 40, 467–490, <https://doi.org/10.1146/annurev-environ-102014-021217>, 2015.
- 5 Giorgi, F. and Mearns, L. O.: Approaches to the simulation of regional climate change: A review, *Reviews of Geophysics*, 29, 191–216, <https://doi.org/10.1029/90RG02636>, <http://dx.doi.org/10.1029/90RG02636>, 1991.
- Giorgi, F., Jones, C., and Asrar, G.: Addressing climate information needs at the regional level: the CORDEX framework, *W.M.O. Buletin*, 58, 175–183, 2009.
- Gultepe, I. and Milbrandt, J. A.: Probabilistic Parameterizations of Visibility Using Observations of Rain Precipitation Rate, Relative Humidity, and Visibility, *Journal of Applied Meteorology and Climatology*, 49, 36–46, <https://doi.org/10.1175/2009JAMC1927.1>, <https://DOI.org/10.1175/2009JAMC1927.1>, 2010.
- 10 Hourdin, F., Musat, I., Bony, S., Braconnot, P., Codron, F., Dufresne, J.-L., Fairhead, L., Filiberti, M.-A., Friedlingstein, P., Grandpeix, J.-Y., Krinner, G., LeVan, P., Li, Z.-X., and Lott, F.: The LMDZ4 general circulation model: climate performance and sensitivity to parametrized physics with emphasis on tropical convection, *Clim. Dyn.*, 27, 787–813, <https://doi.org/10.1007/s00382-006-0158-0>, <http://dx.DOI.org/10.1007/s00382-006-0158-0>, 2006.
- 15 Huang, H.-L., Yang, M.-J., and Sui, C.-H.: Water Budget and Precipitation Efficiency of Typhoon Morakot (2009), *J. Atmos. Sci.*, 71, 112–129, <https://doi.org/10.1175/JAS-D-13-053.1>, 2014.
- Jaeger, E. B. and Seneviratne, S. I.: Impact of soil moisture–atmosphere coupling on European climate extremes and trends in a regional climate model, *Climate Dynamics*, 36, 1919–1939, <https://doi.org/10.1007/s00382-010-0780-8>, <https://doi.org/10.1007/s00382-010-0780-8>, 2011.
- 20 Jourdiere, B.: Ressource éolienne en France métropolitaine : méthodes d'évaluation du potentiel, variabilité et tendances, *Climatologie: École Doctorale Polytechnique*, 2015. Français. ph:+33 01238226, pp. 1–229, 2015.
- Katragkou, E., García-Díez, M., Vautard, R., Sobolowski, S., Zanis, P., Alexandri, G., Cardoso, R. M., Colette, A., Fernández, J., Gobiet, A., Goergen, K., Karacostas, T., Knist, S., Mayer, S., Soares, P. M. M., Pytharoulis, I., Tegoulis, I., Tsikerdekis, A., and Jacob, D.: Regional climate hindcast simulations within EURO-CORDEX: evaluation of a WRF multi-physics ensemble, *Geosci. Model Dev.*, 8, 603–618, <https://doi.org/10.5194/gmd-8-603-2015>, <https://www.geosci-model-dev.net/8/603/2015/>, 2015.
- 25 Knist, S., Goergen, K., Buonomo, E., Christensen, O. B., Colette, A., Cardoso, R. M., Fealy, R., Fernández, J., García-Díez, M., Jacob, D., Kartsios, S., Katragkou, E., Keuler, K., Mayer, S., Meijgaard, E., Nikulin, G., Soares, P. M. M., Sobolowski, S., Szepszo, G., Teichmann, C., Vautard, R., Warrach-Sagi, K., Wulfmeyer, V., and Simmer, C.: Land-atmosphere coupling in EURO-CORDEX evaluation experiments, *Journal of Geophysical Research: Atmospheres*, 122, 79–103, <https://doi.org/10.1002/2016JD025476>, 2014.
- 30 Kotlarski, S., Keuler, K., Christensen, O. B., Colette, A., Déqué, M., Gobiet, A., Goergen, K., Jacob, D., Lüthi, D., van Meijgaard, E., Nikulin, G., Schär, C., Teichmann, C., Vautard, R., Warrach-Sagi, K., and Wulfmeyer, V.: Regional climate modeling on European scales: a joint standard evaluation of the EURO-CORDEX RCM ensemble, *Geoscientific Model Development*, 7, 1297–1333, <https://doi.org/10.5194/gmd-7-1297-2014>, 2017.
- 35 Kunkel, B. A.: Parameterization of Droplet Terminal Velocity and Extinction Coefficient in Fog Models, *Journal of Climate and Applied Meteorology*, 23, 34–41, [https://doi.org/10.1175/1520-0450\(1984\)023<0034:PODTVA>2.0.CO;2](https://doi.org/10.1175/1520-0450(1984)023<0034:PODTVA>2.0.CO;2), [https://DOI.org/10.1175/1520-0450\(1984\)023<0034:PODTVA>2.0.CO;2](https://DOI.org/10.1175/1520-0450(1984)023<0034:PODTVA>2.0.CO;2), 1984.

- Lawrence, D. M., Oleson, K. W., Flanner, M. G., Thornton, P. E., Swenson, S. C., Lawrence, P. J., Zeng, X., Yang, Z.-L., Levis, S., Sakaguchi, K., Bonan, G. B., and Slater, A. G.: Parameterization improvements and functional and structural advances in Version 4 of the Community Land Model, *J. Adv. Model. Earth Syst.*, 3, <https://doi.org/10.1029/2011MS00045>, <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/2011MS00045>, 2011.
- 5 Manabe, S.: Climate and the ocean circulation, 1. the atmospheric circulation and the hydrology of the earth's surface, *Mon. Weather Rev.*, 97, 739–774, [https://doi.org/10.1175/1520-0493\(1969\)097<0739:CATOC>2.3.CO;2](https://doi.org/10.1175/1520-0493(1969)097<0739:CATOC>2.3.CO;2), 1969.
- Mearns, L., Gutowski, W. J., Jones, R., Leung, R., McGinnis, S., Nunes, A., and Qian, Y.: A regional climate change assessment program for North America, *EOS Transactions A.G.U.*, 90, 311–312, <https://doi.org/10.1029/2009EO360002>, 2009.
- Milly, P. C. D.: Potential Evaporation and Soil Moisture in General Circulation Models, *J. Climate*, 5, 209–226, [https://doi.org/10.1175/1520-0442\(1992\)005<0209:PEASMI>2.0.CO;2](https://doi.org/10.1175/1520-0442(1992)005<0209:PEASMI>2.0.CO;2), [https://doi.org/10.1175/1520-0442\(1992\)005<0209:PEASMI>2.0.CO;2](https://doi.org/10.1175/1520-0442(1992)005<0209:PEASMI>2.0.CO;2), 1992.
- 10 Nakanishi, M. and Niino, H.: An Improved Mellor–Yamada Level-3 Model: Its Numerical Stability and Application to a Regional Prediction of Advection Fog, *Bound.-Lay. Meteorol.*, 119, 397–407, <https://doi.org/10.1007/s10546-005-9030-8>, <http://dx.doi.org/10.1007/s10546-005-9030-8>, 2006.
- Nielsen-Gammon, J. W., Powell, C. L., Mahoney, M. J., Angevine, W. M., Senff, C., White, A., Berkowitz, C., Doran, C., and Knupp, K.: Multisensor Estimation of Mixing Heights over a Coastal City, *Journal of Applied Meteorology and Climatology*, 47, 27–43, <https://doi.org/10.1175/2007JAMC1503.1>, <https://DOI.org/10.1175/2007JAMC1503.1>, 2008.
- 15 Nikulin, G., Jones, C., Giorgi, F., Asrar, G., Büchner, M., Cerezo-Mota, R., Christensen, O. B., Déqué, M., Fernández, J., Hänsler, A., van Meijgaard, E., Samuelsson, P., Sylla, M. B., and Sushama, L.: Precipitation Climatology in an Ensemble of CORDEX-Africa Regional Climate Simulations, *Journal of Climate*, 25, 6057–6078, <https://doi.org/10.1175/JCLI-D-11-00375.1>, <https://DOI.org/10.1175/JCLI-D-11-00375.1>, 2012.
- 20 Oleson, K. W., Lawrence, D. M., B. G., Flanner, M. G., Kluzek, E., J. P., Levis, S., Swenson, S. C., Thornton, E., Feddema, J., Heald, C. L., Lamarque, J.-F., Niu, G.-Y., Qian, T., Running, S., Sakaguchi, K., Yang, L., Zeng, X., Zeng, X., and Decker, M.: Technical Description of version 4.0 of the Community Land Model (CLM), NCAR Technical Note NCAR/TN-478+STR, <https://doi.org/10.5065/D6FB50WZ>, 2010.
- 25 Price, C. and Rind, D.: A simple lightning parameterization for calculating global lightning distributions, *J. Geophys. Res.: Atmospheres*, 97, 9919–9933, <https://doi.org/10.1029/92JD00719>, <https://agupubs.onlinelibrary.wiley.com/doi/abs/10.1029/92JD00719>, 1992.
- Ruti, P. M., Somot, S., Giorgi, F., Dubois, C., Flaounas, E., Obermann, A., Dell'Aquila, A., Pisacane, G., Harzallah, A., Lombardi, E., Ahrens, B., Akhtar, N., Alias, A., Arsouze, T., Aznar, R., Bastin, S., Bartholy, J., Béranger, K., Beuvier, J., Bouffies-Cloch e, S., Brauch, J., Cabos, W., Calmanti, S., Calvet, J.-C., Carillo, A., Conte, D., Coppola, E., Djurdjevic, V., Drobinski, P., Elizalde-Arellano, A., Gaertner, M., Gal n, P., Gallardo, C., Gualdi, S., Goncalves, M., Jorba, O., Jord , G., L'Heveder, B., Lebeau-pin-Brossier, C., Li, L., Liguori, G., Lionello, P., Maci s, D., Nabat, P.,  nol, B., Raikovic, B., Ramage, K., Sevault, F., Sannino, G., Struglia, M. V., Sanna, A., Torma, C., and Vervatis, V.: Med-CORDEX Initiative for Mediterranean Climate Studies, *Bulletin of the American Meteorological Society*, 97, 1187–1208, <https://doi.org/10.1175/BAMS-D-14-00176.1>, <https://DOI.org/10.1175/BAMS-D-14-00176.1>, 2016.
- 30 Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Duda, D. M. B. M. G., Huang, X.-Y., Wang, W., and Powers, J. G.: A Description of the Advanced Research WRF Version 3, NCAR TECHNICAL NOTE, 475, NCAR/TN475+STR, 2008.
- Smirnova, T. G., Benjamin, S. G., and Brown, J. M.: Case study verification of RUC/MAPS fog and visibility forecasts, Preprints, 9 th Conference on Aviation, Range, and Aerospace Meteorology, AMS, Orlando, FL, Sep. 2000, 2.3, 6, 2000.

- Stackpole, J. and Cooley, D. S.: Revised method of 1000 mb height computations in the PE model, Technical Procedures Bulletin. U.S. Dept. of Commerce, National Oceanic and Atmospheric Administration, National Weather Service, 57, 6, 1970.
- Vautard, R., Gobiet, A., Jacob, D., Belda, M., Colette, A., Déqué, M., Fernández, J., García-Díez, M., Goergen, K., Güttler, I., Halenka, T., Karacostas, T., Katragkou, E., Keuler, K., Kotlarski, S., Mayer, S., van Meijgaard, E., Nikulin, G., Patarčić, M., Scinocca, J., Sobolowski, S., Suklitsch, M., Teichmann, C., Warrach-Sagi, K., Wulfmeyer, V., and Yiou, P.: The simulation of European heat waves from an ensemble of regional climate models within the EURO-CORDEX project, *Clim. Dyn.*, 41, 2555–2575, <https://doi.org/10.1007/s00382-013-1714-z>, <https://DOI.org/10.1007/s00382-013-1714-z>, 2013.
- WMO: Manual on the Global Observing System, WMO, 544, 2010a.
- WMO: Guide to Meteorological Instruments and Methods of Observation, Weather - Climate - Weather, pp. 1–176, 2010b.
- 10 Wong, J., Barth, M. C., and Noone, D.: Evaluating a lightning parameterization based on cloud-top height for mesoscale numerical model simulations, *Geosci. Model Dev.*, 6, 429–443, <https://doi.org/10.5194/gmd-6-429-2013>, <https://www.geosci-model-dev.net/6/429/2013/>, 2013.
- Yang, M. J., Braun, S. A., and Chen, D.-S.: Water Budget of Typhoon Nari (2001), *Mon. Wather Rev.*, 139, 3809–3828, <https://doi.org/10.1175/MWR-D-10-05090.1>, 2011.
- 15 Yesad, K.: FULL-POS in the Cycle 41T1 of ARPEGE/IFS, Tech. Rep., Meteo-France, <https://www.umr-cnrm.fr/gmapdoc/IMG/pdf/ykfpos43.pdf>, 45, 2015.