Global aerosol modeling with MADE3 (v3.0) in EMAC (based on v2.53): model description and evaluation

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Abstract.

Recently, the aerosol microphysics submodel MADE3 was introduced as a successor to MADE and MADE-in. It includes nine aerosol species and nine lognormal modes to represent aerosol particles of three different mixing states throughout the aerosol size spectrum. Here we describe the implementation of the most recent version of MADE3 into the atmospheric chemistry general circulation model EMAC, including a detailed evaluation of a ten-year aerosol simulation with MADE3 as part of EMAC.

We compare simulation output to station network measurements of near-surface aerosol component mass concentrations, to airborne measurements of aerosol mass mixing ratio and number concentration vertical profiles, to ground-based and airborne measurements of particle size distributions, and to station network and satellite measurements of aerosol optical depth. Furthermore, we describe and apply a new evaluation method, which allows a comparison of model output to size-resolved electron microscopy measurements of particle composition. Although there are indications that fine mode particle deposition may be underestimated by the model, we obtained satisfactory agreement with the observations. Remaining deviations are of similar size as those identified in other global aerosol model studies.

Thus, MADE3 can be considered ready for application within EMAC. Due to its detailed representation of aerosol mixing state, it is especially useful for simulating wet and dry removal of aerosol particles, aerosol-induced formation of cloud droplets and ice crystals as well as aerosol-radiation interactions. Besides studies on these fundamental processes, we also plan to use MADE3 for a reassessment of the climate effects of anthropogenic aerosol perturbations.

1 Introduction

The MESSy (Modular Earth Submodel System; Jöckel et al., 2010) aerosol microphysics submodel MADE3 (Modal Aerosol Dynamics model for Europe, adapted for global applications, 3rd generation) was created with a re-quantification of the aerosol–climate effect of off-shore ship traffic in mind (Kaiser et al., 2014). Simulations in the past indicated that sulfate

5   formed from ship emissions may be one of the major contributors to the negative anthropogenic aerosol radiative forcing (Capaldo et al., 1999; Lauer et al., 2007, 2009; Righi et al., 2011, 2013; Olivié et al., 2012; Peters et al., 2012, 2013). Due to increasingly stringent regulations on the maximum sulfur content of ship fuels (IMO, 2011) this contribution is expected to decrease strongly in the future (Lauer et al., 2009; Righi et al., 2011; Schembari et al., 2012; Johansson et al., 2013; Jonson et al., 2015). Possibly, aerosol nitrate formation will compensate for part of the reduction (Lauer et al., 2009; Bellouin et al., 2011;

10 Righi et al., 2011). Measurements indicated that it may be crucial to take into account coarse mode particle interactions with condensable trace gases in order to quantify this effect (Kerminen et al., 1997; Hara et al., 1999; Yeatman et al., 2001; Cavalli et al., 2004; Nolte et al., 2008; Prabhakar et al., 2014). Such interactions were neglected in the previous assessments, and were therefore included in MADE3, which represents an extension of MADE (Lauer et al., 2007) and its successor MADE-in with enhanced resolution of fine particle mixing state (Aquila et al., 2011).

15 As an improvement to its predecessors, MADE3 includes computationally efficient and consistent representations of three different aerosol mixing states in each of three different size ranges, which can be advantageous for many other applications. For instance, we aim to use MADE3 for assessments of aerosol–ice cloud interactions. Particles composed of compounds with no or very low water solubility (in the following denoted as "insoluble particles"), such as mineral dust or black carbon particles, can serve as ice nuclei initiating ice formation in mixed-phase or cirrus clouds (e.g., Lohmann and Feichter, 2005;

20 Hoose and Möhler, 2012). The ice formation efficiency of these particles strongly depends on their size and state of mixing with soluble aerosol species. To simulate these effects, climate models should allow for explicit predictions of the number concentration, size distribution, and mixing state of aerosol particles containing insoluble components. In the first generation of model studies on the role of ice nuclei in the global climate system, bulk aerosol schemes were applied (Lohmann et al., 2004; Hendricks et al., 2005, 2011). This implied that the number of potential ice nuclei had to be estimated from aerosol mass

25 assuming typical aerosol size distributions. Advanced aerosol schemes allowing the explicit simulation of the aerosol number concentration and size distribution were applied in more recent studies (e.g., Lohmann and Hoose, 2009; Gettelman et al., 2012; Kuebbeler et al., 2014; Zhou and Penner, 2014). However, the individual number concentrations of insoluble particles in different size ranges and mixing states could only partly be quantified with these approaches. MADE3 has the advantage that it allows explicit simulations of the number concentration, size distribution, and mixing state (external or internal mixture) of

30 aerosol particles containing insoluble components. Hence, the new aerosol scheme opens new opportunities for the simulation of aerosol effects on ice clouds.

We intend to use the enhanced capabilities of MADE3 to update the results of our previous studies on the health and climate impacts of the transport sectors (Corbett et al., 2007; Lauer et al., 2009; Winebrake et al., 2009; Righi et al., 2011; Lund et al., 2012; Righi et al., 2013, 2015, 2016). Such model applications will be the subject of future investigations. As a first

step towards these studies, the present article provides a detailed description of the implementation of MADE3 into a global chemistry-climate model as well as an evaluation of a first reference simulation.

The ability of the MADE3 algorithms to solve the gas–aerosol partitioning (outside clouds), new particle formation, and coagulation parts of the aerosol dynamics equation was demonstrated by Kaiser et al. (2014). For the solution of the full
5 equation, a number of further processes have to be considered, namely particle and precursor emissions; particle transport by advection, convection, and diffusion; aerosol precursor chemistry in the gas and liquid phases; cloud and precipitation scavenging of aerosols; as well as their dry deposition and sedimentation. Hence, we describe here the implementation of MADE3 into the atmospheric chemistry general circulation model EMAC (ECHAM/MESSy Atmospheric Chemistry model; Jöckel et al., 2010, 2016), which includes further submodels to represent these processes (Section 2). Subsequently, in Sec. 3 we
10 present an evaluation of the performance of EMAC with MADE3 as a global aerosol model. The evaluation is accomplished by comparison of EMAC aerosol simulations to observational data from a multitude of different sources, including station networks, airborne measurements, laboratory analyses of in situ-sampled particles, and satellite data. The main conclusions of this study are summarized in Sect. 4. Appendix A provides a list of the acronyms used in this article. Details of the aerosol scavenging scheme are explained in Appendix B. Descriptions of the chemistry mechanisms considered are provided in the
15 Supplement.

The work presented in this paper is partly based on the Ph.D. Thesis by J. C. Kaiser (Kaiser, 2016). We therefore signal the reader that significant parts of the text in the Abstract, Sect. 2, Sect. 3 and Appendix B already appeared in Kaiser (2016).

2 Model description

2.1 EMAC setup

20 The EMAC model is a numerical chemistry and climate simulation system that includes submodels describing tropospheric and middle atmospheric processes and their interaction with oceans, land and human influences (Jöckel et al., 2010). It uses the second version of MESSy to link multi-institutional computer codes. The core atmospheric model is the ECHAM5 (5th generation European Centre Hamburg) general circulation model (Roeckner et al., 2006). For the present study we applied EMAC (ECHAM5 version 5.3.02, MESSy version 2.53) in the T42L19 resolution, i.e., with a spherical truncation of T42
25 (corresponding to a quadratic Gaussian grid of approx. 2.8 by 2.8 degrees in latitude and longitude) with 19 vertical hybrid σ-pressure levels up to $10\,\mathrm{hPa}$. The applied model setup comprised the submodels given in Table 1. A model time step length Δt of 30 minutes was used, and a temporal resolution for the simulation output of 12 hours.

We simulated eleven years in "nudged mode", i.e., wind divergence and vorticity, temperature, and logarithm of the surface pressure were relaxed towards ERA-Interim reanalyses for the years 1995–2005. The first simulated year is regarded as the
30 (aerosol) spin-up phase, so that our evaluation only takes into account the ten years that followed.

Emissions of both, gases and aerosol particles, enter the EMAC atmosphere through the submodels OFFEMIS, for prescribed emissions, and ONEMIS, for so-called online emissions that depend on the dynamics of the atmosphere (e.g., wind speed) and the state of its lower boundary (e.g., sea surface temperature). The emission setup used here is described in a separate subsection

Table 1. MESSy submodels used for the present work.

Name	Function	Reference(s)
AEROPT	aerosol optical properties	Dietmüller et al. (2016)
CLOUD	cloud properties and precipitation formation (stratiform clouds)	Roeckner et al. (2006)
CLOUDOPT	cloud optical properties	Dietmüller et al. (2016)
CONVECT	convection parameterizations	Tost et al. (2006b)
CVTRANS	convective transport of trace gases and aerosols	Tost et al. (2010)
DDEP	dry deposition of trace gases and aerosols	Kerkweg et al. (2006a, 2009)
H2O	consistency between water vapor as chemical species and humidity	Jöckel et al. (2006)
JVAL	photolysis rate coefficients for trace gases	Sander et al. (2014)
LNOX	lightning NO_x emissions	Tost et al. (2007)
MADE3	aerosol microphysics	Kaiser et al. (2014)
MECCA	gas phase chemistry	Sander et al. (2011)
OFFEMIS	prescribed emissions	Kerkweg et al. (2006b)
ONEMIS	online emissions	Kerkweg et al. (2006b)
ORBIT	parameters of the earth's orbit around the sun	Dietmüller et al. (2016)
RAD	radiative processes	Dietmüller et al. (2016)
SCAV	cloud and precipitation processing of trace gases and aerosols	Tost et al. (2006a, 2010)
SEDI	aerosol sedimentation	Kerkweg et al. (2006a)
SURFACE	surface properties (e.g., temperature, snow depth, etc.)	Jöckel et al. (2016)
TNUDGE	nudging of trace gas concentrations	Kerkweg et al. (2006b)
TROPOP	tropopause height and planetary boundary layer height	Jöckel et al. (2006)

(Sect. 2.4). Prescribed emissions are representative of the year 2000. The production of nitrogen oxides (NO_x, i.e. NO and NO_2) from lightning was taken into account by the submodel LNOX, using a parameterization by Price and Rind (1992), which is based on convective cloud top height as the driving parameter. The parameterization was tuned to match global total emissions within the observed range (Schumann and Huntrieser, 2007).

5   Aerosol particle transport is part of the tracer advection and vertical diffusion schemes of the base model ECHAM5 and of the convective transport submodel CVTRANS. Horizontal diffusion of particles is not considered in ECHAM5, but it is anyway not expected to contribute significantly to transport on the scales of the model grid boxes as used here.

We used the submodel MECCA to simulate atmospheric gas phase chemistry. For computational efficiency the simplified tropospheric chemistry scheme that was created by Lauer et al. (2007) was used. It includes 34 gases and 60 chemical reactions
10 (47 gas-phase and 13 photolysis reactions) to describe NO_x–HO_x–CH_4–CO–O_3 chemistry and the tropospheric sulfur cycle (see the Supplement for more details). The photolysis rates are calculated by the submodel JVAL. Heterogeneous reactions,

i.e., reactions of trace gases on or with aerosol particle surfaces, are not included. Note, however, that reactions on cloud droplet surfaces are included via the cloud phase chemistry (see Sect. 2.3 and the Supplement).

MADE3, which is used for the representation of aerosol microphysics, will be described in Sect. 2.2. To avoid interpretation of results from feedbacks of the simulated aerosol on model dynamics, we switch off the feedback of the MADE3 aerosol on clouds and radiation in the model configuration described here. Before such feedbacks will be considered, the quality of the MADE3 aerosol has to be proven, which is the purpose of this study.

Deposition of aerosol particles is handled in EMAC by the submodels DDEP, which uses the so-called "big leaf" approach, SEDI, for sedimentation (gravitational settling), and SCAV, for wet deposition. The latter required some MADE3-specific modifications, see Sect. 2.3.

Optical properties of aerosol particles, which are considered to compute aerosol optical depth (AOD) for comparison with satellite data (Sect. 3.5), are determined by the submodel AEROPT. The lookup tables applied in AEROPT are the same for MADE3 and its predecessor MADE. Hence, we used the MADE tables that were created by Lauer et al. (2007) with the help of the software libRadtran (Mayer and Kylling, 2005). Further details on these calculations are provided by Dietmüller et al. (2016).

Cloud properties are calculated by the submodels CLOUD (stratiform clouds) and CONVECT (convective clouds) in EMAC. For stratiform clouds we selected the standard ECHAM5 single-moment cloud scheme (Roeckner et al., 2003), i.e., a scheme that only considers water and ice mass, but no droplet or ice crystal numbers. Although previous studies with the MADE3 predecessors were carried out with two-moment cloud schemes, a single-moment scheme is sufficient here, as we do not attempt to quantify the climatic impact of aerosol particles. This will be the subject of follow-up studies, however. Instead of the Tompkins (2002) method to calculate fractional cloud cover (as described by Roeckner et al., 2003), we choose the parameterization developed by Sundqvist et al. (1989). The high numerical stability of this scheme is advantageous for multi-year climate simulations. For convective clouds we choose the original ECHAM5 scheme (Roeckner et al., 2003), which is based on work by Tiedtke (1989) and Nordeng (1994), with modifications by Brinkop and Sausen (1997).

2.2 MADE3 v3.0

MADE3 was described in detail by Kaiser et al. (2014). Therefore, we only briefly repeat its main characteristics here and in Fig. 1. The aerosol is represented by the modal approach, namely with nine modes that represent different particle mixing states and different particle size ranges. Each of the Aitken, accumulation, and coarse mode size ranges in MADE3 includes three modes: one for particles fully composed of water-soluble species, one for particles mainly composed of insoluble material (i.e., insoluble particles with only very thin coatings of soluble material), and one for mixed particles (i.e., particles composed of soluble material including insoluble immersions). In the following, we will refer to these modes as "soluble", "insoluble", and "mixed" modes, respectively. The considered components that make up these modes are sulfate (SO_4), ammonium (NH_4), nitrate (NO_3), sea spray (SS) components other than chloride (mainly sodium; Na), chloride (Cl), particulate organic matter (POM), black carbon (BC), mineral dust (DU), and aerosol water (H_2O). Different from the MADE3 box model version presented by Kaiser et al. (2014), where the insoluble modes were dominated by BC and mineral dust, we now also

assign hydrophobic POM to the insoluble modes during emission (see below), in order to describe interactions of this aerosol component with clouds more consistently. Observations also show that BC is mostly emitted internally mixed with POM (e.g., Petzold et al., 2013).

MADE3 calculates changes in the particle number concentration, size distribution, and composition due to gas–particle
5   partitioning, particle coagulation, and new particle formation. For the gas–particle partitioning of semi-volatile species, i.e., ammonia (NH_3)/NH_4, nitric acid (HNO_3)/NO_3, and hydrochloric acid (HCl)/Cl, an equilibrium approach is applied, where condensation towards the coarse mode particles is limited to the kinetically possible fluxes. Sulfuric acid (H_2SO_4) and secondary organic aerosol (SOA) precursors are assumed to condense irreversibly on the particles. Besides some technical changes between the MADE3 version described by Kaiser et al. (2014) (v2.0b) and the one used here (v3.0), we modified the treatment
10   of new particles upon nucleation events, as well as the criterion for transferring particles from the insoluble to the mixed modes, denoted as aging criterion, as follows.

The transfer to the mixed modes is induced as soon as insoluble particles obtain a liquid coating of a critical size. We now neglect aerosol water in the aging calculation and, correspondingly, in the target mode assignment upon particle coagulation. Thus we interpret water uptake as a consequence of particle aging rather than as the primary reason. We further neglect the
15   POM fraction in these model operations since its role in the aging process is still uncertain. Particles from the insoluble modes are now transferred to the mixed modes, if the sum of the soluble inorganic component masses exceeds 10 % of the modal dry mass. This assumption is supported by results of laboratory and field measurements as reported by Svenningsson et al. (1994), Khalizov et al. (2009), and Liu et al. (2013). Correspondingly, we assign particles that result from coagulation of insoluble modes with mixed or soluble modes to an insoluble mode in case the resulting soluble inorganic contribution to dry mass is
20   less than 10 %, and to a mixed mode otherwise.

Concerning nucleation we now account for initial growth of particles at unresolved temporal and spatial scales by rescaling the formation rate of H_2SO_4–H_2O particles with a wet diameter of 3.5 nm (as it was formerly implemented) to a formation rate of SO_4 particles with a dry diameter of 10 nm. With this initial size the model seems to perform better at accurately simulating particle evolution, as evidenced by the comparisons of number concentrations and size distributions to observations
25   in the free troposphere, where nucleation is the major source of ultrafine particles (cf. Sects. 3.2 and 3.3). We base this updated assumption on new particle formation measurements as reported by, e.g., Modini et al. (2009), Kerminen et al. (2010), Boulon et al. (2011), Matsui et al. (2011), Young et al. (2013), García et al. (2014), Chandra et al. (2016), Giamarelou et al. (2016) and Ueda et al. (2016).

When coupling the MADE3 aerosol to the cloud processing submodel SCAV (see Appendix B), we assume that the insoluble
30   modes are hydrophobic, whereas we consider mixed and soluble modes to be hydrophilic. Hence, only particles in the mixed and soluble modes undergo liquid nucleation scavenging, i.e., they can serve as nuclei for cloud droplet formation. In contrast, ice nucleation scavenging is considered to be less efficient for purely soluble particles (see Appendix B1). In analogy to Aquila et al. (2011), we assume that 5% of the soluble particles are incorporated into ice crystals (ice nucleation scavenging ratio of 0.05) in cirrus clouds (T≤238.15 K), consistent with scavenging ratios typical for homogeneous freezing of aerosols. In
35   order to account for the ability of insoluble particles to act as ice nuclei in heterogeneous ice formation processes, we assume
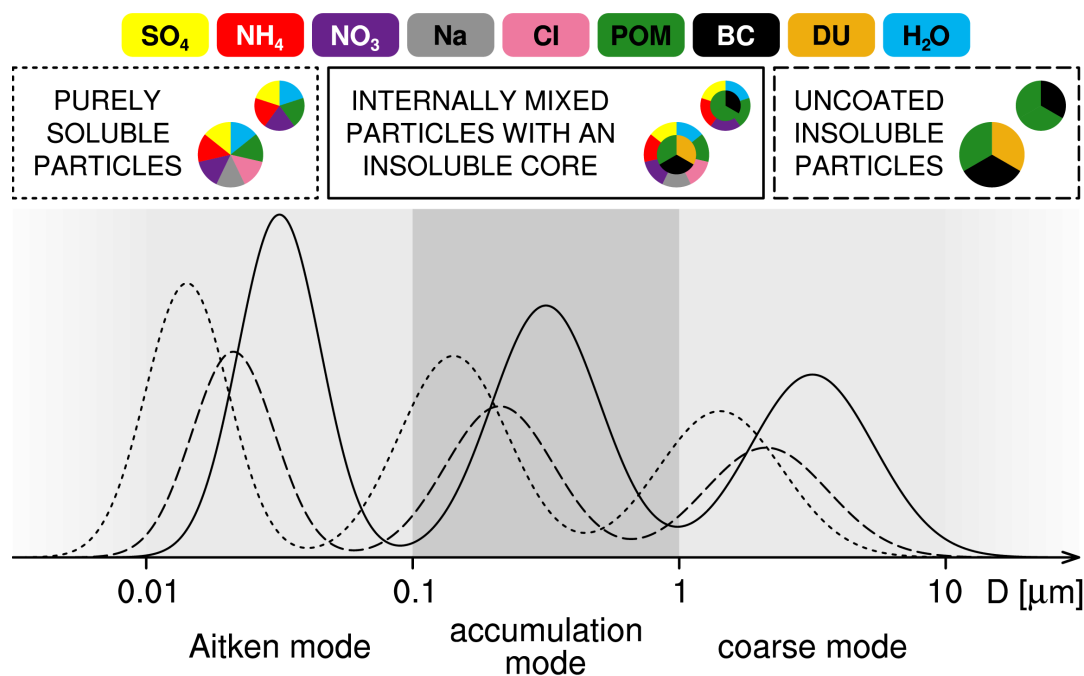
Figure 1. Schematic representation of the MADE3 submodel. The colors represent the different chemical components. The dotted, solid and dashed lines correspond to the different mixing states (soluble, mixed and insoluble, respectively).

a higher scavenging ratio, namely 0.1, for particles containing BC or dust, either externally or internally mixed (insoluble and mixed modes). In the mixed-phase cloud regime (T>238.15 K), we assume an ice nucleation scavenging ratio of 0.1 for all particle types. This rough estimate is based on the fact that only a fraction of cloud droplets freezes during glaciation of liquid clouds via the Bergeron-Findeisen process while the majority of the droplets evaporates releasing large amounts of aerosol mass originally scavenged during liquid droplet formation. A ratio of 0.1 corresponds to typical ratios of the concentrations of ice crystals and cloud droplets in this regime (Korolev et al., 2003). Impaction scavenging does not depend on the particle type.

Due to the extended mode structure of MADE3 with respect to the other two aerosol submodels that can be used with SCAV in EMAC, i.e., the first version of MADE and GMXe (Pringle et al., 2010), a number of modifications to SCAV were required. The main conceptual difference is described in the following subsection.

2.3   Aerosol processing in clouds and precipitation

Cloud and precipitation chemistry as well as wet deposition of both, aerosol and gas species, are treated by the submodel SCAV in EMAC (see Appendix B1). We neglect ice phase chemistry here (incl. uptake of gases onto ice particles), but include 35 chemical components and 45 reactions in the liquid phase chemistry scheme. Heterogeneous processes like the formation of aqueous HNO_3 from gaseous N_2O_5 on droplet surfaces are also considered. We refer the reader to the Supplement for more details on the chemical mechanisms adopted in this work.

Resuspension of aerosol particles from evaporating/sublimating cloud particles or precipitating hydrometeors is an important process to be represented, which has recently been confirmed by Gao et al. (2016). For MADE3 we completely revised the routines describing the redistribution of the resuspended aerosol. The basic assumptions for the assignment of such residual aerosol to the MADE3 modes are described below. Technical details and mathematics are provided in Appendix B2. Note that, in the following, (i) "cloud particles" refers to both, ice crystals and liquid droplets suspended in clouds, as well as to falling snow flakes and rain droplets; (ii) "cloud residual aerosol" also includes "precipitation residual aerosol"; and (iii) "evaporation" includes both, evaporation of cloud and rain droplets, and sublimation or melting plus subsequent evaporation of ice crystals and snow flakes. The following assumptions are made:

1. According to an operator splitting approach, we assume that activation of aerosol particles into cloud particles occurs first, and impaction of interstitial aerosol upon the cloud particles follows in an instantaneous manner as a second step. We acknowledge that this constitutes a strong simplification of the two interdependent processes, which may influence the simulation of the cloud residual aerosol. When more measurement data on the mixing state of cloud residual aerosol becomes available, the impact of this simplification can be evaluated and the algorithm may then be refined. However, as the influence of cloud particle coagulation on mixing state is not represented in the model code, there will be some inevitable error anyway.

2. In order to keep the complexity of the involved equations at a reasonable level and to avoid underestimations of aerosol transformations within clouds, we further assume that, during impaction scavenging, the interactions of interstitial aerosol with cloud particles are as homogeneously distributed as possible over the cloud particle population, regardless of the cloud particles' aerosol cores (see Appendix B2 for more details on this assumption).

3. Aerosol particles from the insoluble modes (dashed line in Fig. 1) cannot be activated into cloud droplets in the present scheme, as they are assumed to be hydrophobic. Nevertheless, they can serve as ice nuclei.

4. The chemical formation of water soluble species within cloud droplets and coagulation between cloud particles lead to accumulation of soluble aerosol components inside cloud particles. To account for such effects we assume that all aerosol particles that were incorporated into cloud particles are hydrophilic upon cloud particle evaporation. Hence, no residual aerosol is assigned to the insoluble, hydrophobic modes.

5. Furthermore, we assume that – due to collection of other aerosol particles, generation of aerosol mass inside cloud droplets, and coagulation of cloud particles – aerosol particle cores of the cloud particles resulting from activated Aitken mode aerosol will have grown from Aitken to accumulation or coarse mode sizes when the cloud particles evaporate. Hence, no residual aerosol is assigned to the Aitken modes.

2.4 Emissions setup

The emissions setup for the present study was in large parts designed by Righi et al. (2013, see their Sect. 2) and includes wind-driven sea spray emissions (Guelle et al., 2001), prescribed emissions of mineral dust and volcanic sulfur (Dentener

et al., 2006), terrestrial dimethyl sulfide (DMS, Spiro et al., 1992), natural SOA precursors (Guenther et al., 1995), as well as prescribed anthropogenic and biomass burning emissions representative of the year 2000 (Lamarque et al., 2010). Emissions of the long-lived greenhouse gases CO_2 and CH_4 are implicitly considered by relaxing their near-surface mixing ratios to observed values for the year 2000, based on data from the Advanced Global Atmospheric Gases Experiment (AGAGE) and the

5 National Oceanic and Atmospheric Administration/Earth System Research Laboratory (NOAA/ESRL). This task is fulfilled by the submodel TNUDGE. For the details on the treatment of aerosol emissions in the model, we refer the reader to Righi et al. (2013). Here, we only describe the parts of the emissions setup that were updated or required MADE3-specific treatment.

Wind speed- and sea surface temperature-dependent oceanic DMS emissions are calculated according to a formulation by Liss and Merlivat (1986) in ONEMIS. In comparison to Righi et al. (2013) we use a more recent climatological near-surface

10 sea water DMS concentration dataset (Lana et al., 2011) as an input to this parameterization.

As aerosol particle number concentrations, size distributions, and mixing states are not included in most of the datasets that we use, we made typical assumptions for mixing states and size distributions here. Following Cooke et al. (1999) and Lohmann et al. (1999), we consider 80 % of the emitted BC and 50 % of the emitted POM to be hydrophobic and assign them to the insoluble MADE3 modes. Note that in the present setup only combustion sources of primary POM are considered and we

15 assume that BC and POM are emitted as internal mixtures in the form of soot particles. Consequently, the hydrophilic fractions are assigned to the mixed modes. Depending on the emission sector, SO_4 is either assigned to soluble or mixed modes. Where nucleation of ultrafine SO_4 particles may play a role even in aged emission plumes, we assign these particles to the soluble Aitken mode. In the other cases we assume that SO_4 is efficiently scavenged by BC/POM particles and, consequently, choose a mixed mode. Volcanic SO_4 is assigned exclusively to the soluble modes, as we do not consider any insoluble particles from

20 volcanic emissions that could bear a coating. Mineral dust emissions are assigned to the insoluble modes, in which particles are assumed to be hydrophobic (Kaaden et al., 2009; Weinzierl et al., 2009), whereas sea spray emissions are assigned exclusively to the soluble modes. Unless explicitly specified in the datasets, i.e., for all emissions except those of sea spray and mineral dust, we derive number emissions from the mass emissions in analogy to the procedure employed by Righi et al. (2013). Under the updated assumptions for the size distribution parameters given in Table 2 the number emissions can thus be computed from

25 the species mass emissions. These number emission fluxes are added to the corresponding MADE3 modes along with the mass emission fluxes from which they were derived.

3 Aerosol evaluation

This section is organized as follows. First, we evaluate simulated near-surface aerosol particle mass concentrations by comparing them to measurements from four different station networks (Sect. 3.1). We then move on to a comparison of the vertical

30 distribution of the simulated aerosol to aircraft measurements of BC mass mixing ratio and of particle number concentration (Sect. 3.2). In Sect. 3.3, we discuss comparisons of simulated size distributions to measurements. Subsequently, we present a method to compare global aerosol model output to size-resolved electron microscopy particle composition measurements, together with a first application (Sect. 3.4). As an aerosol measure derived from the vertical distribution of particle concen-

Table 2. Size distributions assumed for emitted particles. The term $\frac{m}{m_\text{tot}}$ specifies the mass fraction provided by the respective measured or prescribed mode, D_g and σ specify, respectively, the median diameter and geometric standard deviation of the lognormal distributions used to describe the number size distributions of the modes. The "MADE3" columns show the species and modes to which the emissions are assigned. The mode naming convention follows Kaiser et al. (2014): one of the indices "k", "a", "c" is used to specify the Aitken, accumulation, or coarse mode, respectively, and "s", "m", "i" as a second index specifies the soluble, mixed, or insoluble mode, respectively. The value "variable" in the mass fraction and median diameter columns indicates that both, number and mass emissions, are provided by the emission data source, so that number emissions do not have to be derived from the mass emissions. The emission sectors are abbreviated as follows: "AIR" for aviation, "ANT" for anthropogenic non-traffic, "AWB" for agricultural waste burning, "BB" for biomass burning, "LAND" for land transport, and "SHIP" for shipping.

Sector	Mode 1				Mode 2				Reference(s)
	$\frac{m}{m_\text{tot}}$	D_g [nm]	σ	MADE3	$\frac{m}{m_\text{tot}}$	D_g [nm]	σ	MADE3	
Natural emissions									
Sea spray (online)	variable	variable	2.0	Na, Cl: as	variable	variable	2.2	Na, Cl: cs	(Guelle et al., 2001) $+ \sigma_k$ as in MADE3
Mineral dust	variable	variable	1.59	DU: ai	variable	variable	2.0	DU: ci	Dentener et al. (2006)
Volcanic	0.5	30	1.8	SO4: ks	0.5	80	1.8	SO$_4$: as	Dentener et al. (2006)
Anthropogenic and biomass burning emissions									
AIR	0.91	25	1.55	BC: ki, km SO$_4$: ks	0.09	150	1.65	BC: ai, am SO$_4$: am	Petzold et al. (1999)
ANT	—	—	—	—	1.0	138	1.59	BC, POM: ai, am	Birmili et al. (2009)
AWB	1.0	80	1.8	BC, POM: ai, am SO$_4$: am	—	—	—	—	Dentener et al. (2006)
BB	1.0	80	1.8	BC, POM: ai, am SO$_4$: am	—	—	—	—	Dentener et al. (2006)
LAND	0.1 1.0	58 30	1.58 1.8	BC, POM: ki, km SO$_4$: ks	0.9	138	1.59	BC, POM: ai, am	Birmili et al. (2009) Dentener et al. (2006)
SHIP	0.1	70	1.45	BC, POM: ki, km SO$_4$: km	0.9	260	1.25	BC, POM: ai, am SO$_4$: am	Petzold et al. (2008)

trations, composition, and size distributions, we compare simulated AOD to satellite measurements and station network data (Sect. 3.5). Finally, we present global burdens of the simulated aerosol particle species along with the species' tropospheric residence times. Sections 3.1–3.3 and 3.5 draw heavily on model evaluation with the help of the ESMValTool (Eyring et al., 2016).

5   When comparing global model output to observational data, several issues have to be taken into account. As Schutgens et al. (2016a) pointed out, differences of several 10% between simulations and measurements can arise simply due to spatial sampling, when comparing grid box average values to localized observations. Furthermore, observed and simulated microphysical and chemical aerosol properties may not always be fully consistent. Many measurement methods, for instance, are only sensitive to a certain part of the atmospherically relevant particle size spectrum. Specifically, to date, fine mode particles (diameters

10   up to ~ 1–$2.5\,\mu\mathrm{m}$) have received most attention in the literature. This is especially important when comparing "total" aerosol (species) mass and number concentrations. In addition, there is often not a one-to-one correspondence between simulated and measured aerosol species. A prominent issue in this context is related to measurements of '*"black carbon", "soot", "elemental carbon", "equivalent black carbon" and "refractory black carbon"* [which] *synonymously refer to the most refractory and light-absorbing component of carbonaceous combustion particles*' (Petzold et al., 2013). Different and partially inconsistent

15   terminology has been, and is, used in the corresponding literature (Bond and Bergstrom, 2006; Petzold et al., 2013), which has to be kept in mind when comparing simulated BC concentrations to measurement data. Finally, some measurements may be inherently biased due to the method of particle sampling. According to Ames and Malm (2001), for instance, fine mode NO_3 may be underestimated in one station network (CASTNET), while it could be overestimated in another (IMPROVE). The results of the comparisons between MADE3 and the observations, as well as the possible reasons for specific biases, are

20   discussed in the following.

    The discussion here is mostly descriptive and only gives some possible reasons for deviations of simulations from observations. A thorough investigation of such discrepancies would require a large number of sensitivity simulations. Although beyond the scope of the present evaluation, it could be conducted as part of future studies, also for quantification of simulation uncertainties.

25 3.1   Near-surface mass concentrations

Regular measurements within station networks provide both, spatial and temporal data coverage that is well-suited for evaluation of global aerosol models. Biases due to differences in timing of simulation output and observations (Schutgens et al., 2016b) are likely small in this context as the measurements are typically taken by collecting particles on filters over several days and subsequently analysing these samples.

30   Ten-year average simulated near-surface mass concentrations are compared here to the averages of available observational data in the period 1996–2005 from the following station networks: the Interagency Monitoring of Protected Visual Environments network (IMPROVE, Hand et al., 2011) and the Clean Air Status and Trends Network (CASTNET, AMEC Environment & Infrastructure, 2015) in the USA, the European Monitoring and Evaluation Programme network (EMEP, Hjellbrekke, 2014), and the acid deposition monitoring network in East Asia (EANET, Network Center for EANET, 2014). The simulation data

Table 3. Summary of the model–observations statistical comparison of monthly mean near-surface aerosol concentrations from the EMAC simulation with MADE3 and station network data. $\langle OBS \rangle$ ($\langle MOD \rangle$) stands for the arithmetic mean over all data points of the observations (simulation), σ_{obs} (σ_{mod}) for the standard deviations, $fac2$ for the percentage of simulated values that are within a factor of two of the corresponding observational values ($0.5\langle OBS \rangle \leq \langle MOD \rangle \leq 2\langle OBS \rangle$), and N_{pts} for the number of data points, i.e., monthly averages, that went into the comparison. See Appendix A2 in Righi et al. (2013) for more details about the methodology.

Species	$\langle OBS \rangle$ [µg m^{-3}]	σ_{obs} [µg m^{-3}]	$\langle MOD \rangle$ [µg m^{-3}]	σ_{mod} [µg m^{-3}]	$\langle MOD \rangle/\langle OBS \rangle$	$fac2$ [%]	N_{pts}
IMPROVE							
SO$_4$	1.61	1.72	3.09	2.22	1.92	39.7	13162
NH$_4$	1.16	0.56	1.57	0.54	1.36	79.8	609
NO$_3$	0.50	0.75	0.88	0.78	1.77	39.6	13162
POM	1.72	2.01	2.01	1.76	1.17	66.2	13106
BC	0.27	0.27	0.31	0.31	1.14	69.5	13106
CASTNET							
SO$_4$	3.15	2.38	4.22	2.72	1.34	73.2	8215
NH$_4$	1.10	0.77	1.50	0.74	1.36	68.1	8215
NO$_3$	0.87	1.07	1.16	0.90	1.33	45.2	8215
EMEP							
SO$_4$	2.25	1.70	3.10	2.14	1.38	73.2	9365
NH$_4$	0.95	0.90	1.45	0.96	1.52	57.8	3547
NO$_3$	1.68	1.91	1.49	1.25	0.89	53.9	3672
EANET							
SO$_4$	3.36	4.99	3.80	2.48	1.13	57.6	1201
NH$_4$	0.90	1.52	1.50	0.96	1.66	35.3	1181
NO$_3$	0.95	2.09	1.46	1.27	1.53	33.5	1172

is always the sum of the contributions from all modes in the lowermost model layer, i.e., up to $\sim 100\,\mathrm{m}$. The comparison provides an indication of the model's ability to reproduce the climatological state of the northern hemisphere continental aerosol, where anthropogenic emissions are largest. As the emission dataset is assumed to represent year 2000 conditions (Sect. 2.4), the period for the observational data was chosen symmetric to the year 2000. Note, however, that most of the stations (across networks) do not provide complete temporal coverage of the years 1996–2005, which may lead to biases. Specifically, all EANET data that went into the comparison stem from the years after 2000. EMEP and EANET have fewer stations and can thus provide less data than IMPROVE and CASTNET (see Table 3).
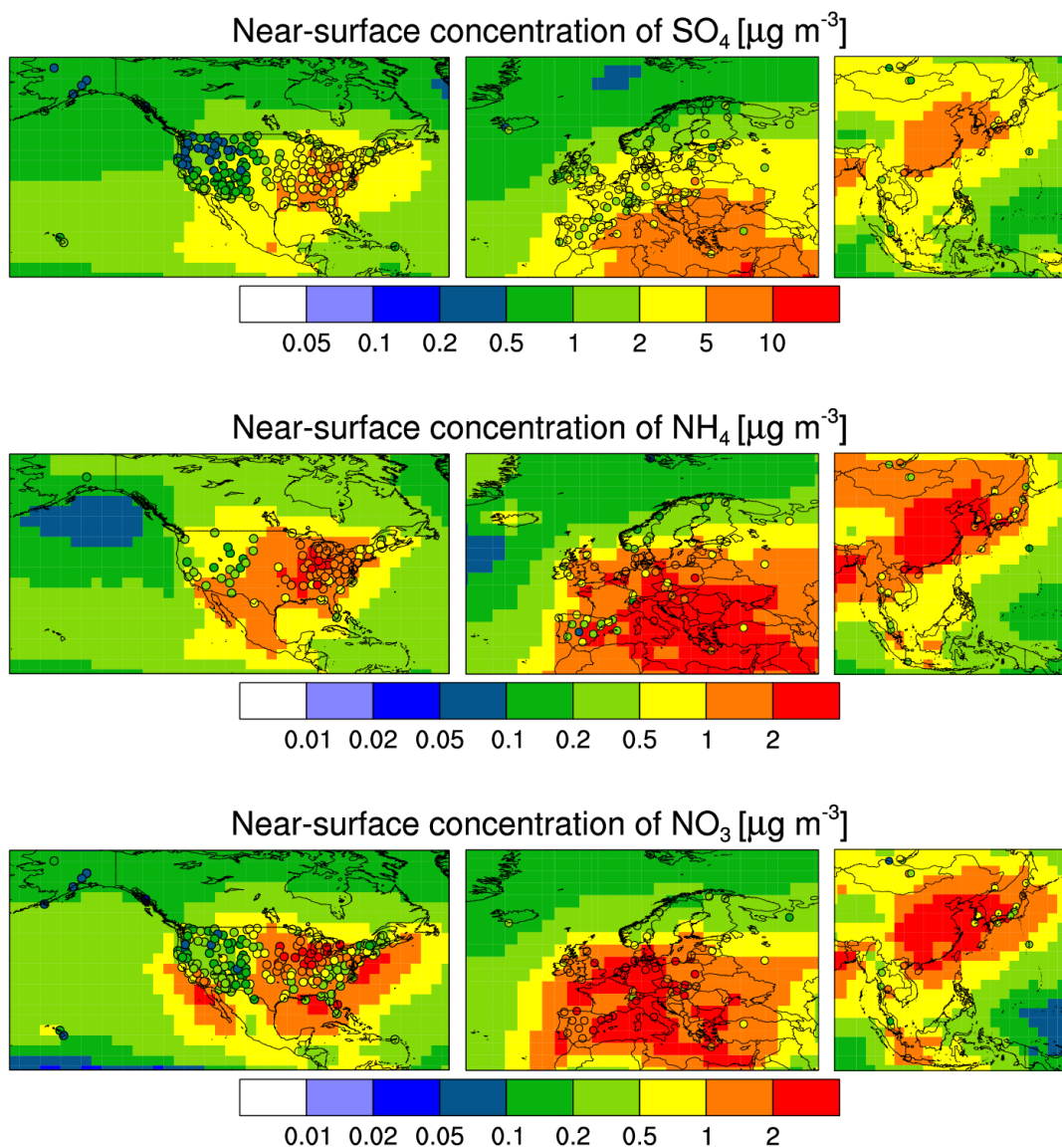
Figure 2. EMAC simulation with MADE3 (background color; "pixels" correspond to the model grid) vs. observations from station networks (filled circles): multi-year average near-surface secondary inorganic aerosol mass concentrations, i.e., SO$_4$ (top), NH$_4$ (middle), and NO$_3$ (bottom). The observational data in the three panels of each row is drawn from CASTNET and IMPROVE (left), EMEP (middle), and EANET (right).

Concentrations of the secondary inorganic aerosol species SO$_4$, NH$_4$, and NO$_3$ are the most widely measured and typically have the longest records, while BC and POM are only measured in the IMPROVE network. This subsection is ordered accordingly: the secondary species are discussed first, the (mostly) primary aerosol components thereafter.

The geographical distribution of near-surface SO_4 concentrations (Fig. 2) is well reproduced over Europe and East Asia, albeit with a small bias (Table 3). Over the USA, agreement between simulated and measured concentrations is better in the east than in the northwest. Hence, the model does not fully capture the west–east gradient seen in the observations. Apart from this discrepancy and the general high bias, however, the spatial SO_4 distribution over the USA is reproduced. The relative

5   deviation of near-surface SO_4 concentrations is notably larger in the case of IMPROVE compared to the other networks (Table 3). As Ames and Malm (2001) do not find systematic differences in SO_4 concentrations between co-located IMPROVE and CASTNET measurements, a possible reason for this difference lies in the locations of the IMPROVE stations. The relative deviations are largest in the northwestern part of the USA (Fig. 2), where most of the stations used in the comparison are IMPROVE stations. These stations are mostly located in national parks and wilderness areas, i.e., in rather clean environments,

10   whereas the large model grid-boxes also cover more polluted areas in the vicinity of these areas.

Observed values of NH_4 concentrations are spatially more heterogeneous than those of SO_4 concentrations, down to scales that cannot be captured by the coarse resolution of the model. Furthermore, emissions of the NH_4 precursor NH_3 are much more uncertain and variable than those of SO_4 precursors. That said, model performance is mostly similar for NH_4 as for SO_4 (Fig. 2). This also means that the west–east gradient over the USA is again underestimated. Note that in case of NH_4,

15   most IMPROVE stations with available data are located in the eastern part of the USA, where agreement of the simulated concentrations with the observations is slightly better than in the west (see also Table 3). The north–south gradient over Europe is generally well captured, with the exception of Spain and the western Mediterranean.

The simulated near-surface NO_3 concentrations agree remarkably well with the observations across Europe (Fig. 2). With respect to the IMPROVE data for the USA it must be noted that several ten percent of the simulated NO_3 mass belong to

20   the coarse modes. In contrast to CASTNET, however, IMPROVE stations only sample particles up to a size of $\sim 2.6\,\mu\text{m}$ (IMPROVE, 1995), so that deviations have to be expected, although partly compensated by the tendency of IMPROVE to overestimate fine mode NO_3. The comparison of NO_3 concentrations to data from CASTNET and EANET yields similar results as for NH_4.

Only IMPROVE stations routinely monitor carbonaceous aerosol components. Hence, the POM and BC simulation vs. ob-

25   servations comparison (Fig. 3) includes only data from the USA. IMPROVE reports mass concentrations for organic carbon (OC), which were converted to POM by multiplying by a factor of 1.4 (Dentener et al., 2006). Both the magnitude and the spatial distribution of simulated near-surface concentrations generally agree well with the observations. This was expected, since the primary aerosol species POM and BC are only marginally affected by uncertainties associated with gas and liquid phase precursor chemistry and gas–particle partitioning which potentially cause distinct discrepancies in the case of secondary

30   species.

To evaluate dust concentrations we follow the same strategy as Aquila et al. (2011), who compared simulated concentrations with the data based on a collection of measurements from 22 stations around the world in the context of the AEROCOM project (Huneeus et al., 2011). The results are shown in Fig. 4: the data from the different stations is grouped in order of increasing dust load according to the respective average dust concentrations (1 to 7: low, 8 to 16: medium, and 17 to 22: high). EMAC

35   (MADE3) generally underestimates dust concentrations, especially in comparison to the medium- and high-concentration
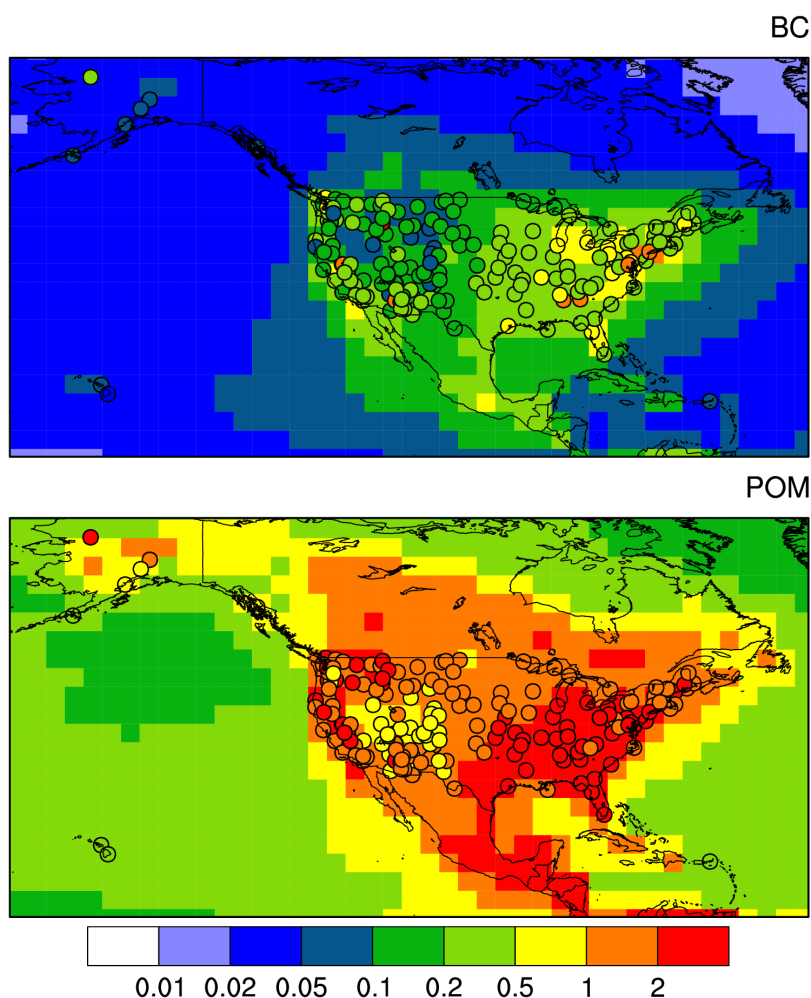
Figure 3. EMAC simulation with MADE3 (background color; "pixels" correspond to the model grid) vs. observations from the station network IMPROVE (filled circles): near-surface BC (top) and POM (bottom) mass concentrations.

stations, whereas the annual cycle is captured reasonably well. As pointed out by Aquila et al. (2011), this discrepancy could be due to the use of an offline monthly-mean climatology for dust emissions in the model, rather than an online, wind-driven dust emission scheme. This could lead to a misrepresentation of atmospheric dust transport and removal. Furthermore, the climatology is representative of the year 2000, which was characterized by relatively low dust emissions. MADE3 simulations

5 with more detailed dust emission parameterizations are planned as a subject of future studies. It is interesting to note, however, that the ability of MADE3 to reproduce dust concentrations has improved considerably over the previous MADE-in version of Aquila et al. (2011), who used the same input climatology for dust emissions.
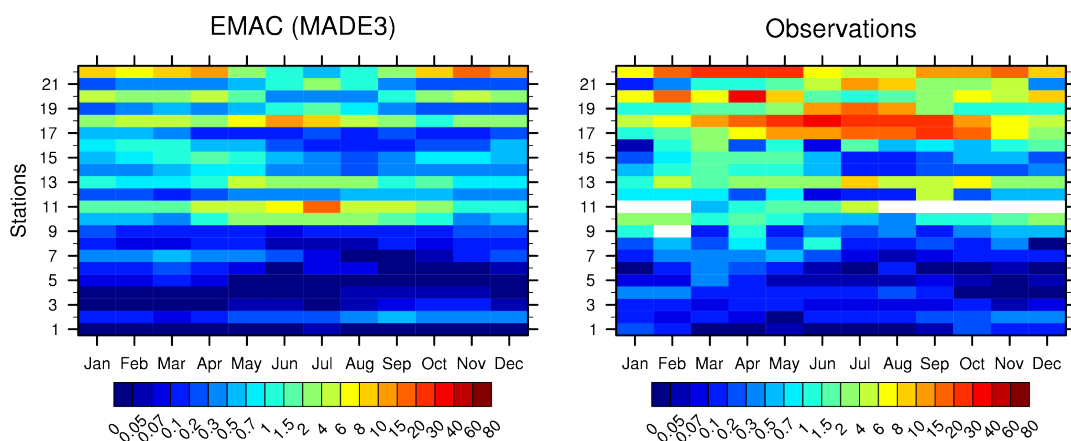
Figure 4. Climatological annual cycle of simulated (left) and measured (right) dust surface-level concentrations in units of µg m^{-3}. The observational data was collected by Huneeus et al. (2011). The location of the 22 stations considered in the plot is shown in Fig. 6 of Aquila et al. (2011).

The station networks also provide measurements of the sea spray components, i.e., Na and Cl. However, their concentrations are extremely low over the continents and, consequently, very sharp land–sea gradients in the near-surface concentrations of these species occur. These gradients cannot be accurately resolved by the model, which complicates the comparison with the observations, especially for stations in costal areas. Hence the station network data is not suited to evaluate sea spray aerosol. For an evaluation of the simulated marine aerosol, we refer to Sect. 3.5 where comparisons with satellite date are presented.

A comparison of simulated near-surface concentrations of various aerosol species to observations at high latitudes is challenging, since observational data for those regions is scarce and only a plausibility check could be performed here. Simulated BC concentrations are close to those measured over two years in North Greenland (as reported by Massling et al., 2015), while simulated SO$_4$ concentrations are roughly a factor of two higher than the measured ones. Total aerosol mass concentrations over Antarctica are larger compared to measurements taken by Hara et al. (2014) during the austral summer 2007-2008.

In conclusion, the simulated near-surface concentrations are mostly larger than the corresponding observed values (Table 3). This high bias is largest for the secondary aerosol components SO$_4$, NH$_4$, and NO$_3$, but only small in the case of BC and POM. It may indicate a too low efficiency of the deposition processes, which might also explain some of the discrepancies discussed in the next subsection that deals with vertical profiles of BC. That the overestimation is generally larger for soluble components indicates that cloud processing may also play a role in the deviations. That said, the statistics of our comparison with IMPROVE data is very similar to those obtained with a previous EMAC version using MADE-in (Aquila et al., 2011). The only exception here is NO$_3$, which could not form on coarse mode particles in MADE-in, and therefore could not reach as high concentrations as in the present study. However, as mentioned above, IMPROVE stations do not provide any insight into NO$_3$ concentrations associated with particles $\gtrsim 2.6\,\mu$m.

The high-bias of near-surface mass concentrations found here is not typically seen in studies with other global aerosol models. Although (relative) discrepancies are often of similar magnitude, the deviations are typically more variable in their directions for different species (e.g., Bauer et al., 2008; Mann et al., 2010; Pozzer et al., 2012; Lee et al., 2015). Nevertheless, the ability to simulate several ten percent of monthly mean values within a factor of two of observations indicates a quality

5 of EMAC with MADE3 that is similar to that of other models (e.g., Pozzer et al., 2012; Kirkevåg et al., 2013). Note that the primary goal of the present study was not to improve on previous aerosol climatologies. Rather, we show that our new model, with its additional capabilities in terms of particle mixing state representation and coarse mode particle interactions, also produces reasonable climatologies and, hence, is ready for investigating new aspects, which could not be addressed with the former versions of the model.

10 **3.2 Vertical distribution**

It is even more delicate to evaluate the global 3-D aerosol distribution than to evaluate the 2-D near-surface distribution. In contrast to the multi-year time series of measurements provided by surface station networks, aircraft measurements only sample aerosol along a specific flight trajectory. Hence, both the spatial and temporal data coverage is limited. Although arguably representative of the season and weather pattern during which flights took place, there is much more uncertainty

15 associated with the comparison of climatological model output to aircraft measurements than with that to station network data. Nevertheless, aircraft campaigns provide a uniquely valuable way to measure vertical aerosol concentration profiles and are routinely used to evaluate the performance of global aerosol models in simulating the vertical aerosol distribution.

Here, we use observational data from campaigns between 1990 and 2014 over the Pacific Ocean, over North and South America, and over Europe, as summarized in Table 4. Both BC mass mixing ratios (aerosol mass per unit mass of air; Fig. 5)

20 as well as aerosol particle number concentrations (Fig. 6) were used. Depending on what a dataset provides, either mean values and standard deviations, or medians and percentiles, or both, are compared between simulation and observations. Where data for individual flights is available, we show the median of each flight in the comparisons. The variability of the measured data includes spatial and temporal concentration variations during and between the flights. The variability of the model output, however, reflects the concentration variations around a climatological state, expressed by long-term mean or median concen-

25 trations, respectively. Simulation output data was selected from the grid boxes that include the flight trajectories and from the output time steps corresponding to the days or months of the year during which the flights took place (Table 4). This means that model data is selected for these days/months for each year of the simulation. Simulated meteorological-induced variability is captured well, even if only data for single days is extracted from each simulated year. For the comparisons we vertically binned both, the simulation, and the measurement data into 1-km intervals. In situ GPS altitude was converted to ambient pressure

30 using U.S. Standard Atmosphere; this enabled the in situ to model comparision.

The general picture that emerges from the comparison of the BC mass mixing ratio profiles (Fig. 5), on the one hand, is a comparatively good resemblance of simulated and observed near-surface mass mixing ratios, particularly in polluted continental regions close to major BC sources. For instance, in case of the DC3 and CONCERT campaigns, the simulated mixing ratios in the planetary boundary layer are close to the observed values. This is consistent with the similarity of the

Table 4. Summary of relevant details and references for the aircraft measurement datasets used in the evaluation of vertical aerosol profiles simulated with MADE3 in EMAC. The values in parentheses in the column "Time" are the numbers of measurement flights considered for the evaluation.

Name	Location	Time (#)	Parameter	Reference
UCN-Pacific	Pacific Ocean	1990-05 (15) , 1995-11 (33), 1996-09 (21), 1999-03 (19)	Particle number	Clarke and Kapustin (2002)
INCA (Punta Arenas) INCA (Prestwick)	Chile Scotland	2000-03/04[*] 2000-09/10[*]	Particle number	Minikin et al. (2003)
Oct-AVE	Texas	2004-11-10 (1), 2004-11-12 (1)	BC mass	Schwarz et al. (2006)
CR-AVE	Costa Rica	2006-02 (3)	BC mass	Schwarz et al. (2008)
TC4	Costa Rica/Panama	2007-08 (5)	BC mass	Spackman et al. (2011)
HIPPO	Pacific Ocean	2009-01[†], 2009-11[†], 2010-03/04[†], 2011-06[†], 2011-08[†]	BC mass	Wofsy et al. (2012) Schwarz et al. (2013)
CONCERT	Central Europe	2011-09 (6)	BC mass	Dahlkötter et al. (2014)
DC3	Central USA	2012-05/06 (12)	BC mass	Huntrieser et al. (2016) Schwarz et al. (2017)
ACCESS	Scandinavia	2012-07 (11)	BC mass	Roiger et al. (2015) Schwarz et al. (2017)
SALTRACE (East)	Cape Verde	2013-06 (5)	BC mass	Weinzierl et al. (2017) Schwarz et al. (2017)
SALTRACE (West)	Eastern Carribean	2013-06/07 (13)	BC mass	Weinzierl et al. (2017) Schwarz et al. (2017)
ACCESS-2	Central California	2014-05 (6)	Particle number	Moore et al. (2017)

[*] 9–10 flights per campaign, numbers not given separately

[†] more than 700 profiles in total, numbers not given separately, number of flights not given
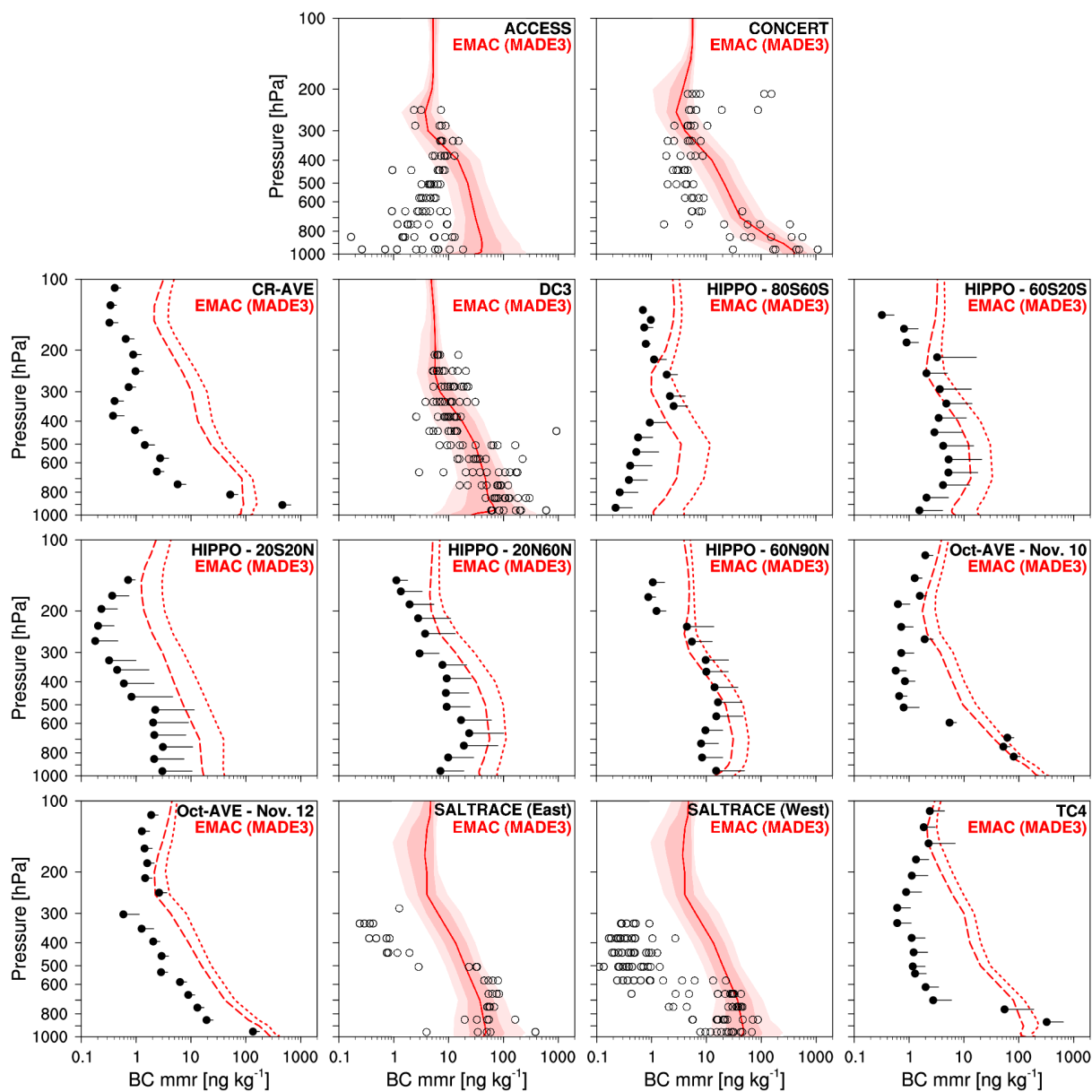
Figure 5. BC mass mixing ratios (mmr) in the EMAC simulation with MADE3 (red) vs. measurements during various field campaigns (black). Dashed lines and filled circles represent mean values, dotted lines and whiskers represent standard deviations, which are only shown in the direction of larger values for clarity. Solid lines stand for median values. Light and dark shadings indicate the 10th to 90th, and 25th to 75th percentiles, respectively. Hollow circles are the median values of individual flights. Descriptions of the campaigns are provided in Table 4 and in the text. Note that the vertical axis of the left plot in each row applies to the other plots of that row as well, and the horizontal axes of the plots in the lowermost row also apply to the plots in the other rows.

IMPROVE station measurements and the corresponding simulation results. On the other hand, the simulated near-surface mixing ratios in remote areas (e.g., the Pacific Ocean covered by the HIPPO campaigns), as well as those simulated for higher altitudes, mostly exceed the corresponding observations. A high bias of up to a factor of 10 occurs in many cases. In this context, it should be mentioned that the data obtained in missions initiated before 2003 used filter-based absorption measurements to

5   infer effective BC concentrations, whereas missions initiated after 2003 used Single Particle Soot Photometers (SP2) to report refractory BC concentration. The BC measurements with SP2 cover a particle size range of about 90-600 nm volume-equivalent diameter, assuming $1.8\,\mathrm{g\,cm^{-3}}$ void-free density, but for some datasets they have been slightly corrected (generally by < 15%) to reflect the total accumulation-mode BC mass. Except for fresh emissions very close to the sources, which most of the data does not represent, we expect the SP2 to capture most of the aerosol BC mass (Schwarz et al., 2006). Although underestimations

10   of the total BC mass in any non-accumulation size modes due to the detection size limits in the measurements cannot be ruled out, discrepancies between model and observation of the order of a factor of 10, as found here, are probably almost certainly insensitive to this much smaller issue.

The high bias of the simulation with respect to the measured profiles could indicate an underestimation of aerosol scavenging as also hypothesized in the previous subsection. In addition, overestimated upward transport, possibly in convective plumes,

15   could also contribute. Ouwersloot et al. (2015) found increased mixing ratios of an artificial tracer in the free troposphere when using an improved convective transport scheme that was recently developed for future versions of MESSy. This may mean that the general tendency of the simulated aerosol mass mixing ratios to exceed the observed values could actually be even larger, i.e. using a similar transport scheme here might lead to even larger discrepancies. Previous studies with MADE (MADE-in) rather showed a negative (slightly negative) bias of simulated vs. measured concentrations (Lauer, 2004; Aquila, 2009), which

20   could indicate that the overestimation in the present work is caused outside the aerosol microphysics submodel, possibly by the representations of scavenging and vertical transport. However, the scavenging efficiency also depends on the aerosol size distribution (see Sect. 3.3), which is largely controlled by the aerosol microphysics submodel. It remains for future studies to investigate the presented hypotheses in detail.

Several other model studies included comparisons to the observational datasets used here. For instance, Lohmann et al.

25   (2007, ECHAM5-HAM) achieved close agreement of the BC mass mixing ratio profiles for the Oct-AVE data, but, using the same model with some modifications to aerosol-related mixed-phase cloud processes, Lohmann and Hoose (2009) found a similar overestimation of the CR-AVE data as that in Fig. 5. While Bauer et al. (2008, MATRIX) could better reproduce the decline of the BC mixing ratios with altitude close to the ground in the CR-AVE and TC4 data, EMAC with MADE3 performs better towards the top of the model domain. In a recent study, Lund et al. (2017) demonstrated that discrepancies between

30   BC simulations with the OsloCTM2-M7 model and the HIPPO data can be strongly reduced by modifications in the model representation of BC wet scavenging. This again demonstrates that deficiencies in the model descriptions of wet removal can play a key role in this context. BC concentrations in the free troposphere are overestimated by many other models (e.g., Koch et al., 2009, 2010; Schwarz et al., 2013; Allen and Landuyt, 2014; Schwarz et al., 2017). Several authors, among them Kipling et al. (2013, HadGEM3-UKCA), Wang et al. (2013, CAM5), and Allen and Landuyt (2014, CAM5), found a better agreement

35   with measured vertical profiles when improving the representation of aerosol–convection interactions. This includes aerosol
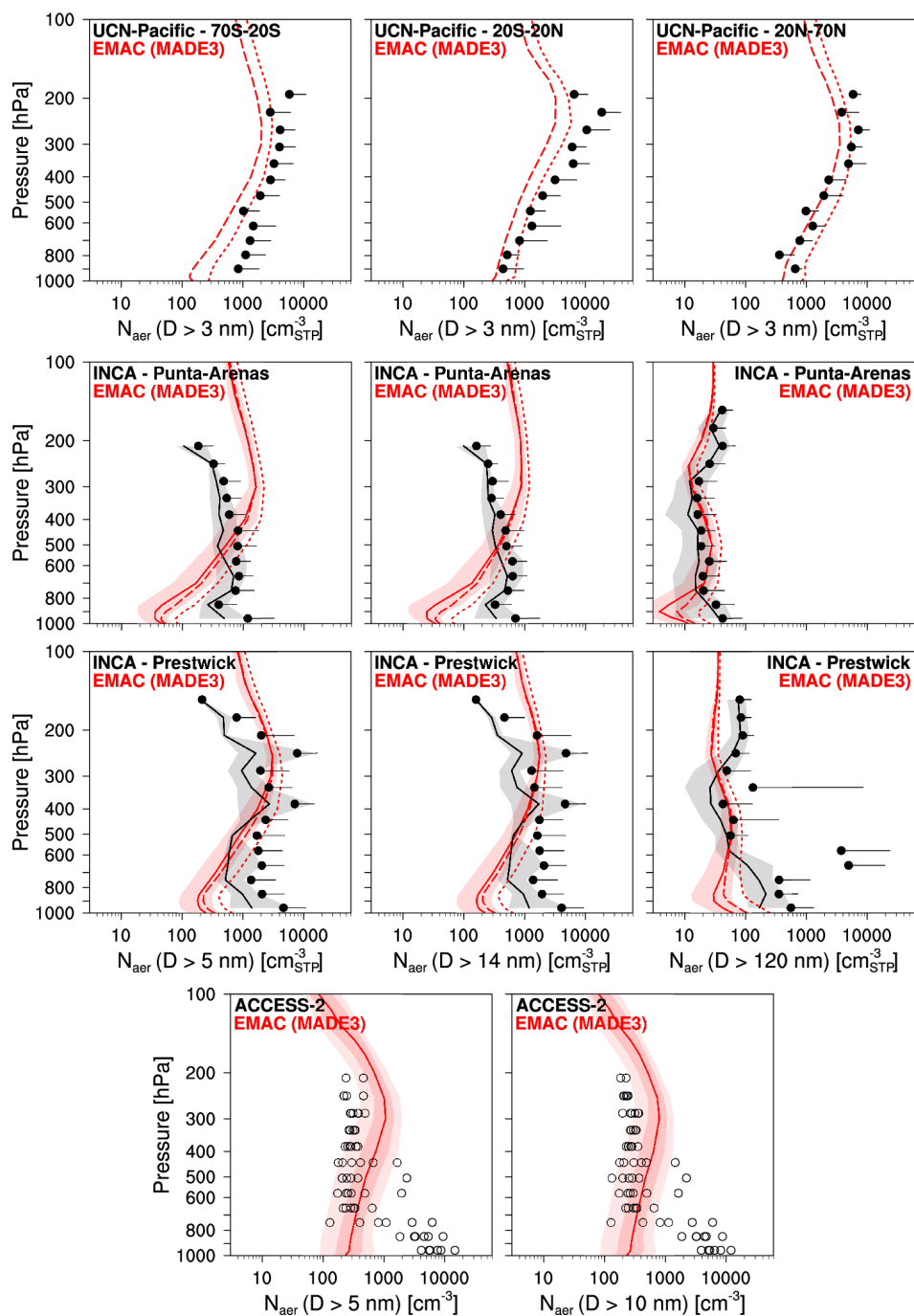
Figure 6. Same as Fig. 5, but for aerosol particle number concentrations (converted to standard temperature and pressure, STP), with various cut-off diameters.

activation, vertical transport, and wet removal in convective clouds. Note, however, that EMAC with MADE3 performs better in the upper tropospheric BC comparison to HIPPO data than the multi-model average of the models that took part in the AEROCOM project (Schwarz et al., 2013).

When comparing simulated aerosol particle number concentration profiles with measurements (Fig. 6) we find a comparatively good agreement over the Pacific Ocean, where both, spatial and temporal coverage, by the observations are most extensive (more than 200 profiles of the UCN Pacific campaign; Clarke and Kapustin, 2002). Note that the simulation values are the result of an integration of the number size distribution from the cut-off diameter (3 nm) upwards. Especially in the northern hemisphere the agreement is remarkable. In the equatorial latitudes, number concentrations agree well in the lower troposphere, but simulated number concentrations are smaller at high altitudes which could be a result of an underestimated efficiency of new particle formation. Over the Southern Pacific aerosol number concentrations are significantly underestimated in the lower troposphere which could be an indication that especially natural sources of aerosol number are underrepresented in the model. For instance, new particle formation mechanisms including natural organic compounds (e.g., Kirkby et al., 2016; Tröstl et al., 2016) are neglected. Since natural precursors might be very relevant for new particle formation in the southern hemisphere where anthropogenic influences are comparatively small, this model deficiency might lead to particularly large discrepencies. The low bias of simulated aerosol number is not in contradiction to the high bias of the BC concentrations discussed above, since aerosol number is controlled by the large concentrations of ultrafine particles, which provide only very small contributions to aerosol mass and which are distinctively smaller than BC particles.

Similar to the comparison with the southern hemispheric UCN-Pacific data, the simulated lower tropospheric aerosol number concentrations are smaller than observed during INCA and ACCESS-2, which again could be a consequence of missing aerosol sources in the model. This deficiency can even affect the concentration of larger aerosol particles in the cloud condensation nuclei size range as reflected by the comparison with the INCA data. Hence future work should focus on improving the representation of natural background aerosol, as also concluded by several other global aerosol modelling studies (e.g., Carslaw et al., 2013, 2017). In some cases, particularly in comparison to the INCA campaign, the model shows higher ultrafine particle number concentrations in the upper troposphere. A possible reason could be an overestimated nucleation rate. Zhang et al. (2012, ECHAM-HAM2) obtained a strong reduction in nucleation mode number concentrations between $\sim 400\,\mathrm{hPa}$ and $\sim 150\,\mathrm{hPa}$ when switching from the Vehkamäki et al. (2002, 2013) scheme employed in MADE3 to a more recent parameterization. This shall motivate future studies with MADE3 to investigate the role of the nucleation scheme in more detail.

3.3 Size distributions

Size distributions provide more detailed information on the aerosol population than integral particle number concentrations. Unfortunately, however, suitable observational data for our evaluation is scarce, especially when it comes to measurements above the ground. For the present study we compared simulated size distributions to data from ground-based measurements (Putaud et al., 2003; Van Dingenen et al., 2004) and aircraft-based observations (Petzold et al., 2002). The latter have the particular advantage that size distributions were determined for different altitudes throughout the troposphere. Simulated size distributions are taken from the grid boxes corresponding to the geographical coordinates of the measurements, and only from
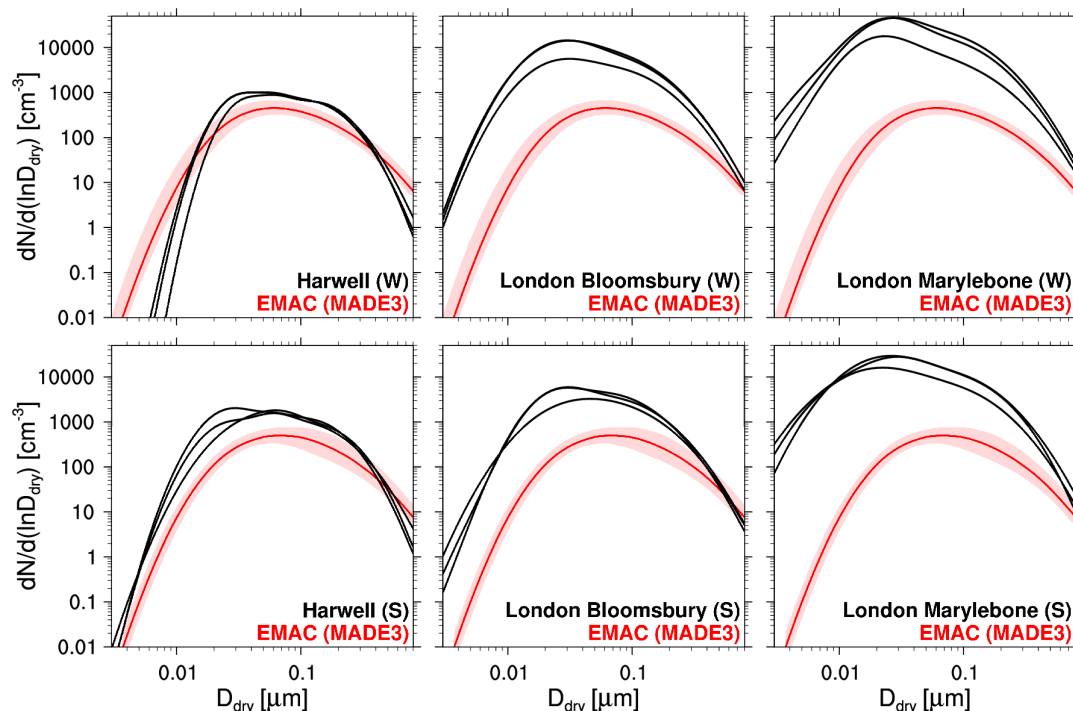
Figure 7. Aerosol particle number size distributions in the EMAC simulation with MADE3 (red) vs. ground-based measurements (Putaud et al., 2003; Van Dingenen et al., 2004) during winter (W, top row) and summer (S, bottom row) at the same locations (columns). Each plot contains three measured size distributions: one for the morning, one for the afternoon, and one for the night hours. The three stations represent the following conditions (left to right): rural, urban, and kerb-side (terminology adapted from Putaud et al., 2003). All measurement locations fall into the same model grid box, so that the simulated size distribution only differs between the top and bottom rows, but not between columns. Solid lines stand for median values, shadings indicate the 25th to 75th percentiles. Note that the vertical axis of the left plot in each row applies to the other plots of that row as well, and the horizontal axes of the plots in the lowermost row also apply to the other plots in the respective columns.

those time steps (in each simulated year) that correspond to the days or months of the observations. We found that variability due to model meteorology is captured well with this approach.

The ground-based measurements were performed at ten European stations that monitored aerosol particle size distribution during at least one full season, i.e., either winter (December, January, February) or summer (June, July, August) in the 1990s or early 2000s. Putaud et al. (2003) fitted up to three log-normal modes to the measured distributions for three times of the day, namely the morning, the afternoon, and the night. Figure 7 shows a subset of our comparisons, which serves to illustrate our results and the problems associated with this type of evaluation.

Our main conclusion here is that the comparability of simulated and measured size distributions can be strongly affected by the specific characteristics of the local environments at the respective stations. This is especially evident when comparing simulation output to data from three stations that fall into the same model grid box, as we do in Fig. 7. Only one size dis-

tribution can be realized in this grid box at any given time in the model. While the model agrees comparatively well with the measurements at the rural station (left column), it shows distinctively smaller concentrations for all particle sizes when compared to measurements from the urban background station (middle column) and the kerb-side station (right column). This was expected, since local concentration enhancements occurring close to local sources cannot be resolved by the model with

5 its large horizontal resolution of about 300 km. The rural station (Harwell) might occasionally be influenced by urban pollution since it is located in the vicinity of London. Hence, the long-term median concentration at this station is expected to be closer to the large-scale median concentration of the model grid box than the median values from the other two stations, as the grid box contains both urban and rural environments. This interpretation is consistent with comparisons to measurements at natural background stations (other grid boxes, not shown) where the model shows larger average concentrations than the observations

10 which are expected to be representative of the less polluted fractions of the respective grid boxes.

Simulated near-surface size distributions over Europe appear to be strongly affected by the emissions. Hence, as most (prescribed) emissions in our simulation are considered as monthly averages, we see little variability of the size distributions. Another deficit of the MADE3 aerosol particle size distribution is the MADE3 output appearing almost unimodal in many cases, whereas the observations often show two or more distinct modes. This finding is consistent with the result of the box

15 model test of MADE3 (Kaiser et al., 2014), and we now find it to be independent of season and location. The discrepancy may be caused partially by differences between the size distributions assumed to calculate particle number emission rates and the size distributions obtained after assigning these emissions to the respective MADE3 modes (Sect. 2.4). In most cases the widths of the modes that were fit to the measured data are narrower than those assumed in MADE3, where $\sigma = 1.7$ and $\sigma = 2.0$ for the Aitken and accumulation modes, respectively. Simulations with alternative assumptions on mode widths are intended

20 to be the subject of future studies.

Further possible contributions to deviations between the simulation and the observations could be related to the timing of simulation output and measurements as well as to the new particle formation approach employed in MADE3. The lack of temporal collocation of simulation output with measurement times may already bias our results (Schutgens et al., 2016b). In addition, Lee et al. (2013a) and Lee et al. (2013b) found that boundary layer nucleation of new particles could contribute up

25 to several 10 % to the uncertainty in number concentrations of particles larger than 50 nm. MADE3 includes an empirical nucleation scheme (Vehkamäki et al., 2002, 2013). As several model studies (e.g., Spracklen et al., 2006; Matsui et al., 2013; Makkonen et al., 2014; Pietikäinen et al., 2014) suggest, the incorporation of more advanced nucleation schemes can lead to a more accurate reproduction of observed aerosol particle number concentrations. Global simulations with other new particle formation schemes could be the subject of a future study with MADE3 in EMAC.

30 Despite the mentioned shortcomings, EMAC performs reasonably well with MADE3 compared to the published results of other global model–observation comparisons of aerosol particle size distributions. Neither Wang et al. (2009, CAM3-IMPACT) nor Zhang et al. (2012, ECHAM5-HAM2) achieved substantially closer agreement with their models. Even Lee et al. (2015, ModelE2-TOMAS) found similar deviations despite using a sectional model, which would be expected to resolve size distributions better than modal models with fixed mode widths.
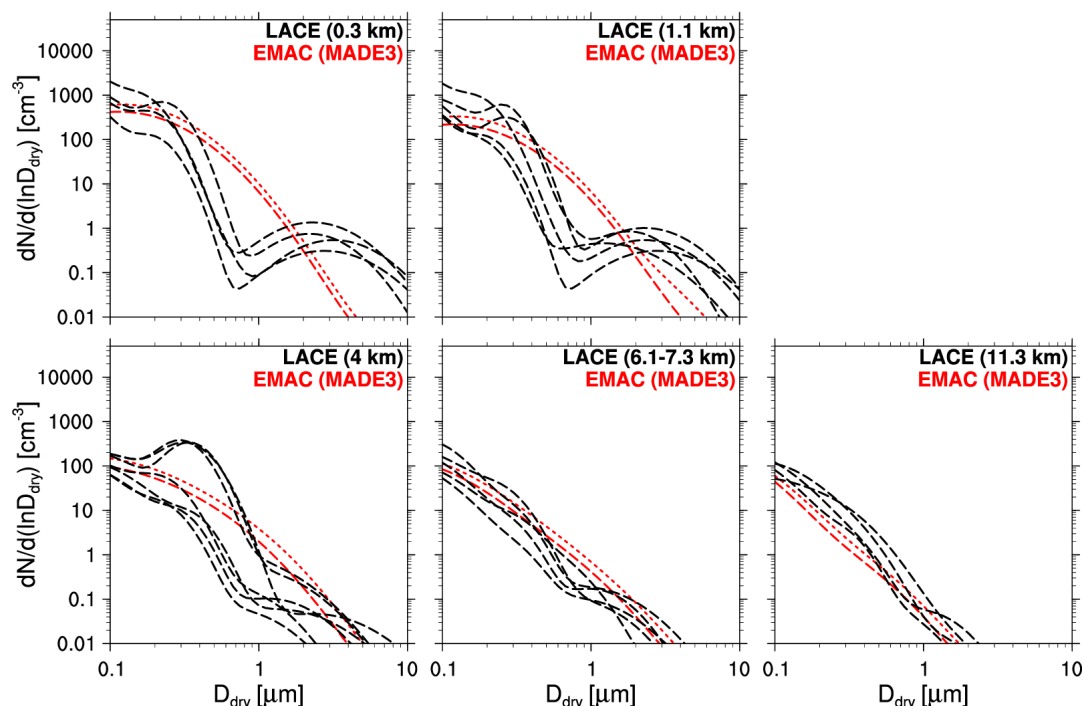
Figure 8. Aerosol particle size distributions in the reference simulation with MADE3 (red) vs. measurements (black) for four to five individual flights (depending on altitude) during the LACE campaign over northeastern Germany. Dashed lines represent mean values, dotted lines represent standard deviations, which are only shown in the direction of larger values for clarity.

Many of the arguments presented above also apply to the comparison of our simulation data to data from the LACE campaign (Fig. 8). The measurements were taken during July and August 1998 at different altitudes over northeastern Germany (Petzold et al., 2002). We use three-mode fits to the measured size distributions for four to five individual flights here, depending on the flight altitude. Again, the fitted modes are much narrower ($\sigma \leq 1.6$) than the MADE3 modes.

5 Notable differences between simulation and LACE data include the lack of the coarse mode at lower altitudes and the accumulation mode peak in the upper boundary layer/lower free troposphere from the simulation output. As the measured coarse mode declines with altitude, it may have to do with local, anthropogenically-induced dust emissions that are not included in the emission dataset used here. The peak at $\sim 300\,\mathrm{nm}$ in the upper boundary layer/lower free troposphere measurements was caused by a forest fire aerosol layer that cannot be reproduced in the simulation because this specific fire is not contained

10 in the emission dataset.

Looking at the remaining parts of the size spectrum and considering the model's capabilities, we see good agreement of the simulated size distributions with the LACE data. Furthermore, we find that agreement improves with altitude, i.e., with increasing particle age.

3.4 Size-resolved composition

To enable a specific evaluation of the new coarse mode particle representation in MADE3 it is useful to compare model output to size-resolved particle composition measurements. However, such data rarely includes coarse mode particles and the correspondence between simulated and measured quantities is not always straightforward. We therefore present a strategy for

5 evaluating simulated size-resolved aerosol composition with the help of electron microscopy data of in-situ sampled aerosol particles. For an initial application of this strategy we chose a dataset from measurements performed in January and February 2008 at a ground station at Praia, Cape Verde, by Kandler et al. (2011) during the SAMUM-2 field campaign (Ansmann et al., 2011).

Particle sizes as determined in the electron microscopy measurements are given as equivalent circle diameters of the par-

10 ticles' projected areas. We assume that these can be directly compared to the diameters derived from the simulated aerosol particle number and component mass concentrations, the mode widths, and the assumed component densities, since spherical particles are assumed in the model (Kaiser et al., 2014).

The experimental analysis is performed on individual particles, i.e., 48599 particles in the dataset used here. Thirteen major elements were detected in the investigated particle population. Based on the relative contributions of the elements to the particle

15 volume, each particle is assigned to one of 12 different particle classes, e.g., sulfates, chlorides, oxides, silicates (see Kandler et al., 2011, for details).

For the comparison to model output this procedure has a severe drawback. It would require classification of the MADE3 particles according to the same, or analogous, rules as the measured particles. However, since all particles within each MADE3 aerosol mode are assumed to have the same composition (model assumption of perfect internal mixture of all involved com-

20 pounds), only particles of a maximum of nine different compositions can coexist at the same time in each grid box of the model. Classification of model particles, or rather modes, is therefore not reliable from a statistical point of view. For instance, consider a mode that contains both, SO_4 and Cl. With the model assumption of perfect internal mixtures, its total volume can always be assigned to only one class, either to the sulfates or to the chlorides. In reality, however, the mode would likely contain both particles with a major contribution from sulfate (assigned to the sulfates class by the measurements) and particles with

25 a major contribution from chloride (assigned to the chlorides class by the measurements). Hence, classification of the model modes would create unacceptable sampling biases.

Furthermore, nitrogen compounds only produce weak signals in the measurements, and material from the sampling substrates can affect the analysis of carbonaceous matter. Of the species simulated by MADE3, only SO_4, Na, Cl, and DU can therefore be determined reliably in the measurements of contributions to particle composition.

30 For these reasons we adopted a different view on the electron microscopy data. In the approach employed here, the component masses of each analyzed particle are assigned to one of five diameter "bins" according to the particle's size. Only those components that can be compared to model output are considered. The chloride fraction is measured directly, and considered to derive exclusively from sea spray. The sodium fraction is also measured directly, but for correspondence to the MADE3 Na tracer (which represents the whole non-chloride sea spray mass in the model), the sea spray sulfate fraction has to be added.
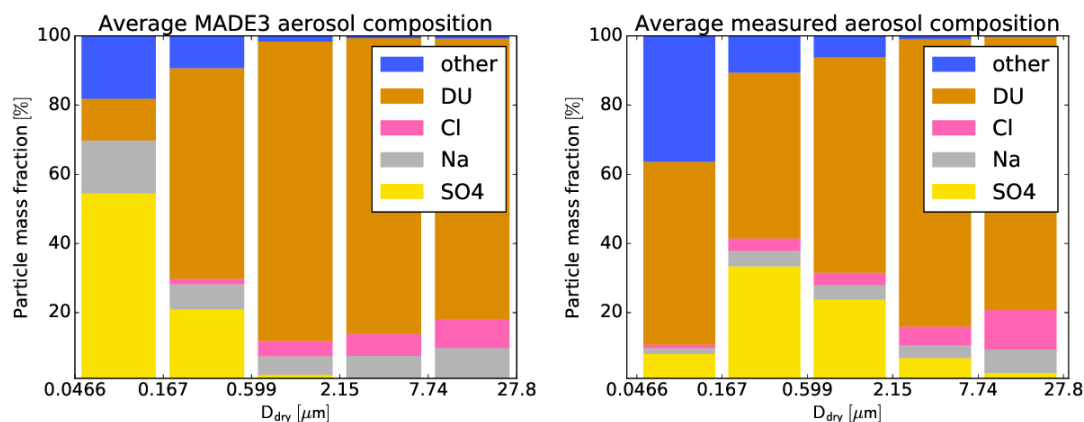
Figure 9. Average size-resolved aerosol composition as simulated by the model (left) and as measured during the SAMUM-2 campaign (right). Only the mass fractions of species that can be compared between measurement data and model output are depicted.

The latter can be derived from the measured chloride fraction under the assumption of a typical sea spray composition, i.e., 54.6 % of chlorine atoms, and 2.82 % of sulfur atoms, and under the assumption that all this sulfur is present in the form of sulfate. The rest of the detected sulfur is also assumed to stem from sulfate and can be compared to the MADE3 SO_4 tracer. The mineral dust contribution is derived from multiple elements that are typical of mineral dust (Silicon, Aluminum, Iron,

5   Magnesium, Calcium, Potassium, Phosphorus, Titanium, and Sodium).

Model output is binned into the same diameter intervals as the measurement data by integrating the mass size distribution of each mode from the lower to the upper bin boundary and then summing up the contributions of the individual modes. Thus, measurement data and model output are brought to the same format and can be compared.

An example comparison is shown in Fig. 9. The measurement panel (right) displays the average particle composition over

10   the whole SAMUM-2 campaign (26 individual days, or 48205 particles). The rest of the analyzed particles fell outside the size range presented here. With 3729 particles, the rightmost bin has the smallest data base. For the model plot 12-hourly output from the grid box that contains the measurement station was averaged over the ten evaluated simulation years, considering only the days of the year when the measurements took place.

The preliminary result shown in Fig. 9 must be interpreted with caution. It is not possible to exactly reproduce the conditions

15   during the SAMUM-2 campaign with the model setup used here, except by chance. Especially, the monthly mean year 2000 DU and SO_4 emissions in the simulation may not be representative of the actual situation in the beginning of the year 2008. Moreover, the meteorological features of the simulated years (1996-2005), which largely impact the simulated mean aerosol properties at the measurement site, might not correspond well to the specific meteorological conditions in 2008. Local pollution sources cannot be resolved by the model either. That said, the comparison reveals similarities between the simulated and

20   measured data in the decrease of the SO_4 fraction and the increases of the sea spray (Na plus Cl) and DU fractions with increasing size. Major discrepancies, however, exist in the composition of the smallest compared particles. We also analyzed the model biases in the individual years, but the interannual variability (not shown) was found to be small, hence meteorology

alone cannot explain the discrepancies. On the other hand, the SEM analysis in particular of the smallest size fraction might have a bias towards an underestimation of sulfate particles due to their instability under the electron beam. A thorough analysis, including also comparisons of the measured and simulated size distributions, but also measurement uncertainties, should be the subject of a separate study.

5   ## 3.5   Aerosol optical depth

AOD provides an integral measure of the vertical aerosol column. On the one hand, it can be computed from the simulated aerosol properties discussed in the previous subsections, i.e., particle composition, particle sizes, and their vertical distributions. On the other hand, AOD can also be derived from measurements with ground-based and satellite-borne radiometers. Here (Fig. 10) we compare the simulated AOD to data from the ground-based AErosol RObotic NETwork (AERONET, Holben

10   et al., 1998, 2001), and against satellite data from the ESACCI Swansea University (SU) ATSR-2 v4.21 aerosol product (North et al., 1999; Bevan et al., 2012; Holzer-Popp et al., 2013; de Leeuw et al., 2015) and from the MODerate resolution Imaging Spectroradiometer (MODIS) Level 3 Collection 6 data (Levy et al., 2013). Since annual mean AOD regionally changed by up to 10 % during the last decade (2000–2010, e.g., Yoon et al., 2014; Pozzer et al., 2015), only year 2000 data is used in the comparisons of annual mean AOD here. This does not apply to MODIS, for which we considered the year 2003, i.e. the earliest

15   year available in the time period covered by the instrument.

The results shown in Fig. 10 reveal that, in comparison to the measurements, the model simulates up to ∼50% higher AOD in the major pollution and biomass burning plumes that originate in East Asia, central Africa, and South America. This high bias is consistent with the general tendency of the model to overestimate aerosol mass concentrations as seen in Sects. 3.1 and 3.2. These high mass concentrations may also entail a higher aerosol water content, which could increase AOD further. The

20   sensitivity of AOD to differences in hygroscopic growth, i.e., water uptake due to aerosol particle hygroscopicity, was recently demonstrated, e.g., by Li et al. (2014). Furthermore, the simulated AOD can be very sensitive to the scavenging scheme for aerosol particles entrained into convective clouds. The choice of the scheme can lead to several 10 % different annual mean AOD values (Croft et al., 2012). A better agreement is found in case of less polluted areas such as the remote oceans. The sign of deviations occurring in these areas shows spatial variations. This could be an indication for misrepresentations of either

25   natural aerosol sources or long-range transport of anthropogenic particles and should be the subject of further investigations in future studies. The model tends to underestimate AOD where DU is abundant, especially over the Sahara and the Arabian peninsula. Potential reasons for this underestimation include the use of prescribed monthly mean year 2000 DU emissions, the assumption on the DU size distribution upon emission, and the DU representation in AEROPT, the submodel that computes aerosol optical properties (Sect. 2.1). Johnson et al. (2012) and Nabat et al. (2012) found improved agreement of simulated

30   AOD with observations when using a parameterization with more of the emitted DU mass in the coarse mode. Furthermore, the studies by Zhao et al. (2013) and Mahowald et al. (2014) indicated that a modal representation of DU particles with fixed mode widths may have unavoidable shortcomings.

AOD data from different satellite instruments does not agree perfectly, but shows similar patterns in large parts of the globe. Consistent patterns of the two satellite-model differences may therefore indicate areas where the model could be biased. Re-
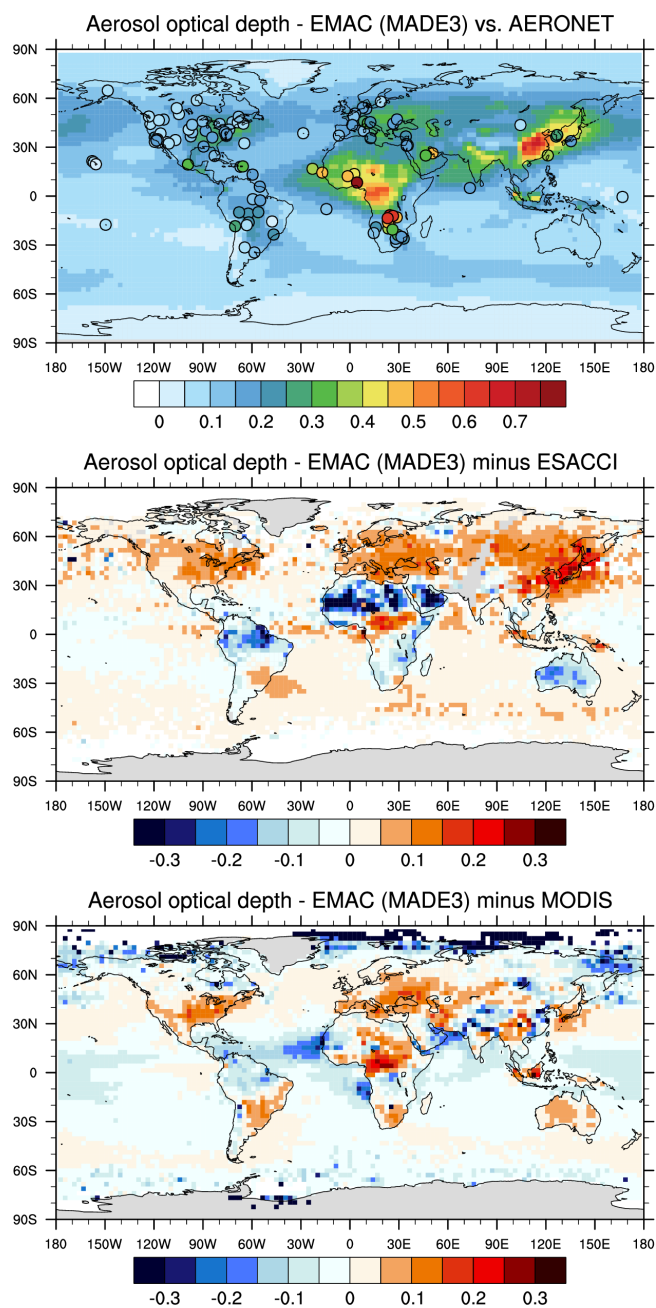
Figure 10. Annual mean AOD in the reference simulation with MADE3 (top, background color) vs. observations from the AERONET network (top, filled circles) and vs. satellite data from the ESACCI Swansea University (SU) ATSR-2 v4.21 aerosol product (middle) and from MODIS Level 3 Collection 6 (bottom). "Pixels" in the panels correspond to the model grid. Results are shown for the year 2000 (AERONET and ESACCI) and for 2003 (MODIS).

maining satellite-satellite differences do even occur for data from the same instruments, if it is obtained with different retrieval algorithms (e.g., Popp et al., 2016). Hence, one cannot expect perfect agreement of simulated AOD with the observations either. It has for example been shown that MODIS AOD is larger by about 0.03 on global average than ESACCI SU values, while the ESACCI SU algorithm shows significantly overestimates AOD in dust regions such as the Sahara (Lauer et al., 2017). Further-
5 more, uncertainties involved in the model calculations of particle optical properties can of course also contribute to deviations. While different models have different strengths and weaknesses, it is interesting to note that many models have a low bias in AOD on a global annual average basis (e.g., Pozzer et al., 2012; Kirkevåg et al., 2013; van Noije et al., 2014; Lee et al., 2015; Michou et al., 2015) rather than a tendency towards a high bias as seen here: 18% (-5%) with respect to the ESACCI (MODIS) data. Relative underestimations in some of the mentioned studies are actually larger than these values, so that we can claim
10 reasonable performance of EMAC with MADE3 as a global aerosol model. Deviations of the order of $10\,\%$ should not be over-interpreted anyway, as the lack of temporal collocation of the simulation output with the measurement times can already lead to biases of this magnitude (Schutgens et al., 2016b).

3.6 Tropospheric burdens and residence times

Although it is not an evaluation in the sense of a check against observational data, a comparison of global tropospheric aerosol
15 burdens and residence times to estimates from other model studies is also instructive. The burden $m_{\mathrm{tot},a}$ of aerosol species a is computed here as the sum over the volume integrals of the mass concentrations c_a in all grid boxes. The species' residence times, $t_{\mathrm{res},a}$, are derived from the burdens and the sums of the deposition fluxes, $F_{\mathrm{dep},a}$, as

$$t_{\mathrm{res},a} = \frac{m_{\mathrm{tot},a}}{F_{\mathrm{dep},a}}. \tag{1}$$

Simulated burdens and residence times are presented in Table 5 along with results from other modeling studies.
20   The values simulated with MADE3 in EMAC mostly fall within the ranges of previous estimates. For the secondary inorganic species, i.e., SO_4, NH_4, and NO_3, the partitioning between the coarse and fine modes appears to play an important role. While SO_4 and NH_4 are found predominantly in the fine modes ($> 95\,\%$ on average), NO_3 partitions roughly equally between the fine and coarse modes on average. While the SO_4 and NH_4 burdens simulated in the present study are at the upper end of the range of available model results, the NO_3 burden lies well within the range spanned by the other studies. This could be an indication
25 that especially the life times of fine mode aerosol are comparatively long in our model. The comparatively high NH_4 burden in our simulation should be interpreted with care, since corresponding values have been reported by only three other studies.

4   Summary, conclusions and outlook

We implemented the aerosol microhpysics submodel MADE3 into the global chemistry-climate model EMAC as a successor to MADE and MADE-in. The new submodel version includes nine aerosol species and represents three types of aerosol particles
30 in three different size ranges. As a first application, we performed a ten-year model simulation. To evaluate the model quality, we compared the simulation output to data from a wide range of observations. These include aerosol (species) mass and

Table 5. Annual average values for global tropospheric aerosol burdens and residence times simulated with MADE3 in EMAC. The ranges of values for the ten evaluated years are given in brackets. Furthermore, the ranges of values from a compilation of other studies (Lauer and Hendricks, 2006; Textor et al., 2006; Bauer et al., 2008; Wang et al., 2009; Mann et al., 2010; Pozzer et al., 2012; Kirkevåg et al., 2013; He and Zhang, 2014; van Noije et al., 2014; Lee et al., 2015; Michou et al., 2015) are shown in the table. Note that this is not meant to be a comprehensive literature review, but should yield a representative picture. Extreme values may stem from specific sensitivity studies.

| Species | Burden [Tg] | | Residence time [days] | |
| | This work | Other studies | This work | Other studies |
	Mean [Range]	Range	Mean [Range]	Range
SO4	3.16 [2.95; 3.22]	[0.90; 3.35]	4.75 [4.61; 4.84]	[1.28; 5.60]
NH4	0.88 [0.83; 0.90]	[0.1; 0.46]*	5.22 [5.09; 5.31]	[1.49; 4.50]*
NO3	0.55 [0.54; 0.56]	[0.02; 1.64]	1.65 [1.63; 1.68]	[1.37; 6.40]
Na	2.24 [2.18; 2.31]	[1.95; 3.04]	0.24 [0.23; 0.25]	[0.19; 1.20]†
Cl	1.72 [1.68; 1.79]	[4.47; 4.52]‡	0.18 [0.18; 0.19]	–
POM	1.68 [1.62; 1.73]	[0.99; 4.79]	4.32 [4.23; 4.42]	[2.56; 9.52]
BC	0.19 [0.19; 0.20]	[0.08; 0.34]	4.07 [3.99; 4.13]	[2.44; 9.60]
DU	9.20 [9.00; 9.40]	[8.63; 41.6]	1.76 [1.71; 1.81]	[1.56; 5.92]

* only reported by Lauer and Hendricks (2006), Pozzer et al. (2012) and He and Zhang (2014).

† values for sea spray.

‡ only reported by He and Zhang (2014).

number concentrations, size distributions, and AOD from surface-based, airborne, and satellite measurements. The results of these comparisons are summarized below.

The main conclusion from the near-surface mass concentration comparisons is that EMAC with MADE3 mostly captures the observed annual average spatial patterns of all aerosol species included in the model. Best agreement was obtained for black

5 carbon and particulate organic matter, but they could only be compared over the USA. Among the other species, quantitative agreement is typically best for SO_4, with up to $\sim 70\%$ of the simulated monthly mean values within a factor of two of the observations (factors between 0.5 and 2). Concentrations of the nitrogen-containing components, NH_4 and NO_3, are spatially less heterogeneous in the simulations than in the observations. This is likely caused by the coarse model resolution and by higher temporal variability of the precursor emissions compared to those of SO_4, which leads to larger uncertainties in the

10 emission datasets. We detected a high bias of the average of the simulated values vs. the observations for nearly all species, which might have to do with underestimated removal of fine mode particles from the atmosphere. Note, however, that near-surface mass concentrations could only be evaluated over the northern hemisphere continents, with very few exceptions. The comparisons demonstrate the ability of EMAC with MADE3 to simulate several ten percent of the monthly mean aerosol species concentrations within a factor of two of the observations, which indicates a quality of the model that is similar to that

15 of other global aerosol models.

The comparison of vertical BC mass mixing ratio and aerosol particle number concentration profiles revealed that the model representations of aerosol vertical transport and wet removal may need to be improved in order to avoid overestimations of the upper tropospheric aerosol load. In addition, the model description of new particle formation needs to be further developed towards more robust representations of particle formation from inorganic and also organic aerosol precursors. In this context, it should be stressed again that discrepancies in the representation of vertical aerosol profiles are a common feature of current global aerosol models and need to be the subject of in-depth investigations and resulting model improvements in the future.

Simulated near-surface size distributions, or rather their level of agreement with observations, was strongly affected by the coarse spatial and temporal model resolution. The simulated distributions agreed well with measurements in areas representative of continental background conditions. However, a rather unimodal shape could often be seen in the simulation, whereas observed distributions contained a separate nucleation mode in many cases, for instance. This could be due to the coarse model resolution which impedes the representation of local enhancements of ultrafine particles due to local emissions. It could also be a consequence of the relatively wide MADE3 modes in comparison to those fitted to the observational data. Furthermore, weaker seasonal variability was found in the simulation than in observations across Europe. For the future, we plan to deepen this analysis by extending the simulated period and include comparisons to the data collected by Asmi et al. (2011) and Birmili et al. (2016). This shoud also include simulations with alternative assumptions for the mode widths.

The comparison of simulated AOD to ground-based and satellite observations provided further evidence for some of the conclusions drawn above. Compared to the observational data, our model shows larger AOD in regions affected by anthropogenic pollution and biomass burning emissions. In contrast, the simulated AOD is smaller compared to the observations over regions where DU dominates the aerosol composition. Together with deviations in sea-spray-dominated areas, this shows the necessity of improving the respresentation of wind-driven dust and sea spray emissions in the model.

Our evaluation also included a comparison with electron microscopy measurements, suggesting that the model is largely capable to simulate the dependence of aerosol composition on particle size. However, these analyses need to be extended in the future to draw more robust conclusions. Simulations of specific episodes during which measurements were taken are required for this purpose, with the appropriate meteorology and emissions. The size distribution of the measured aerosol particles should also be taken into account for a thorough comparison, especially that of the surface area available for SO_4 condensation, in order to understand the model–observation differences in the representation of fine mode sulfate in more detail. Such data, however, has to be measured with different instruments. For some campaigns both, size-resolved composition and size distribution measurements, are available. MADE3 can be evaluated with these data in the future applying the method presented here.

We mentioned many sources of uncertainty in the parameters and parameterizations that are part of the aerosol microphysics and transport calculations, e.g., related to new particle formation and convective transport or scavenging. Besides, there are numerous issues that have to be taken into account when comparing simulations to observations. Among those are the specific meteorological conditions and emissions, which influenced the measured aerosol properties, the correspondence of measured and simulated species, and the uncertainties inherent in the observations, which are rarely reported. A detailed analysis of all these factors is beyond the scope of the present study. Our main conclusion here is that, in all "disciplines", the simulation with MADE3 achieved a level of agreement with observations that falls within the range of results reported by other authors from

simulations with their models. The same mostly holds for burdens and residence times of the MADE3 aerosol components, so that the new submodel can be considered ready for application.

One of the intended applications of MADE3 in EMAC is the re-assessment of the aerosol-induced ship emissions effect on climate as described in the Introduction. We saw much higher aerosol nitrate concentrations over the major shipping routes
5   in our present simulation than in previous simulations with MADE, where interactions of the coarse mode with the gas phase were limited to the exchange of water in the condensed and gaseous phases. Hence, previous conclusions, especially in terms of the assessment of low-sulfur fuel scenarios, might have to be reconsidered.

Furthermore, MADE3 will be used as part of EMAC to assess climate effects of the aerosol through modification of ice and mixed-phase cloud properties. MADE3 is especially suitable for such applications due to its mixing state representation with
10 fully soluble, mixed, and insoluble particles in each of the three size ranges of the Aitken, accumulation, and coarse modes.

5   Code availability

MESSy is continuously further developed and applied by a consortium of institutions. The usage of MESSy, including MADE3, and access to the source code is licensed to all affiliates of institutions which are members of the MESSy Consortium. Institutions can become members of the MESSy Consortium by signing the MESSy Memorandum of Understanding. More informa-
15 tion can be found on the MESSy Consortium Website (http://www.messy-interface.org). The model configuration discussed in this paper has been developed based on version 2.53 and will be part of the next EMAC release (version 2.54).

The Earth System Model eValuation Tool (ESMValTool), used for evaluating MADE3 in comparison with observations, is available under doi:10.17874/ac8548f0315 (see also www.esmvaltool.org).

6   Data availability

20 The model simulation data analyzed in this work is available upon request and will be published via doi together with the final version of this manuscript.

Appendix A: List of acronyms

ACCESS	Arctic Climate Change, Economy, and Society project
ACCESS-2	Alternative fuel effects on Contrails and Cruise EmiSSions 2
AEROCOM	AEROsol Comparisons between Observations and Models
AERONET	AErosol RObotic NETwork
AGAGE	Advanced Global Atmospheric Gases Experiment
AOD	Aerosol Optical Depth
ATSR	Along Track Scanning Radiometer
BC	Black Carbon
CAM	Community Atmosphere Model
CASTNET	Clean Air Status and Trends NETwork
Cl	Chloride
CONCERT	CONtrail and Cirrus ExpeRimenT
CR-AVE	Costa Rica Aura Validation Experiment
DC3	Deep Convective Clouds and Chemistry project
DMS	DiMethyl Sulfide
DU	Mineral DUst aerosol
EANET	Acid Deposition Monitoring Network in East Asia
ECHAM5	5th generation European Centre Hamburg
ECMWF	European Centre for Medium-Range Weather Forecasts
EMEP	European Monitoring and Evaluation Programme
ESMValTool	Earth System Model eValuation Tool
EMAC	ECHAM/MESSy Atmospheric Chemistry model
ESACCI	European Space Agency Climate Change Initiative
GMXe	Global Modal-aerosol eXtension
HadGEM3	Hadley Global Environmental Model 3
HAM	Hamburg Aerosol Model
HIPPO	HIAPER Pole-to-Pole Observations
IMPACT	Integrated Massively Parallel Atmospheric Chemical Transport
IMPROVE	Interagency Monitoring of Protected Visual Environments network
INCA	INterhemispheric differences in Cirrus properties from Anthropogenic emissions
LACE	Lindenberg Aerosol Characterization Experiment
MADE	Modal Aerosol Dynamics model for Europe, adapted for global applications
MATRIX	Multiconfiguration Aerosol TRacker of mIXing state
MESSy	Modular Earth Submodel System
MODIS	MODerate resolution Imaging Spectroradiometer
Na	Sea spray components other than chloride (mainly sodium; Na)
NH_4	Aerosol ammonium
NO_3	Aerosol nitrate
NOAA/ESRL	National Oceanic and Atmospheric Administration/Earth System Research Laboratory
Oct-AVE	October 2004 Aura Validation Experiment
POM	Particulate Organic Matter
SALTRACE	Saharan Aerosol Long-range Transport and Aerosol-Cloud-Interaction Experiment
SAMUM-2	SAharan Mineral dUst experiMent 2
SO_4	Aerosol sulfate
SOA	Secondary Organic Aerosol
SP2	Single Particle Soot Photometer
SS	Sea Spray aerosol
TC4	Tropical Chemistry, Cloud, and Climate Coupling
TOMAS	TwO-Moment Aerosol Sectional microphysics model
UCN	Ultrafine Condensation Nuclei
UKCA	United Kingdom Chemistry and Aerosols model

Appendix B: SCAV: Technical aspects

The submodel SCAV has undergone numerous updates since its original publication (Tost et al., 2006a), e.g., concerning the ice phase as described by Tost et al. (2010). Here, we briefly summarize the most important parts of the algorithms as currently implemented for MADE3 aerosol, as not all of them have been documented in the literature so far.

B1 Aerosol scavenging in and below clouds

Four operators are applied in the following sequence:

1. ice nucleation scavenging,

2. liquid nucleation scavenging,

3. snow impaction scavenging,

4. rain impaction scavenging.

In SCAV "nucleation scavenging" refers to both the actual nucleation of cloud droplets or ice crystals, and the scavenging of aerosol by cloud particles due to the aerosol particles' Brownian motion. The latter is currently not included in many global aerosol models, although it may have a substantial impact on particle number concentration (Pierce et al., 2015). The term "impaction scavenging" summarizes the processes through which aerosol particles are taken up by precipitation, i.e., falling hydrometeors. These processes include Brownian motion of aerosol particles towards hydrometeors, impaction of hydrometeors upon aerosol particles, and interception of aerosol particles by hydrometeors. A scavenging rate, η, which is applied to the number and mass concentrations of aerosol particles in grid boxes with clouds and/or precipitation, is computed for each mode and each of the above operators. It represents the aerosol fraction of the respective mode incorporated into cloud or precipitation particles during a model time step Δt.

Constant scavenging rates are assumed for ice nucleation:

$$
\eta_{\mathrm{inu}} = \begin{cases} \frac{0.1}{\Delta t} & \text{for } T > 238.15\,\mathrm{K} \\ \frac{0.1}{\Delta t} & \text{for } T \leq 238.15\,\mathrm{K} \quad \text{in MADE3 modes with insoluble cores} \\ \frac{0.05}{\Delta t} & \text{for } T \leq 238.15\,\mathrm{K} \quad \text{in purely soluble MADE3 modes,} \end{cases}
\tag{B1}
$$

where T is absolute temperature.

Cloud droplet nucleation is taken into account via an empirical function (Tost et al., 2006a) that is applied to the hydrophilic particles, i.e., to those in the soluble and mixed MADE3 modes:

$$
\eta_{\mathrm{lnu}}^{\mathrm{nuc}} = \frac{2}{\pi \Delta t} \arctan\left(\left[2.5\,\tilde{D}_{\mathrm{g}}\right]^6\right),
\tag{B2}
$$

with the dimensionless number median diameter of the respective aerosol mode $\tilde{D}_{\mathrm{g}}$ (in units µm, as for D_{g}). For Brownian motion a semi-empirical formulation of the scavenging coefficient by Pruppacher and Klett (1997) is used, which leads to

a scavenging rate of (Tost et al., 2006a)

$$\eta_{\text{lnu}}^{\text{Br}} = \frac{1}{\Delta t}\left[1 - \exp\left(-\frac{1.35\,\text{LWC}\,\Delta_p}{r_{\text{cld}}^2}\Delta t\right)\right].$$ (B3)

Here, LWC is the cloud liquid water content (mass per unit volume), and r_{cld} is the effective cloud droplet radius, which is set constant at $r_{\text{cld}} = 17.5 \times 10^{-6}$ m. The aerosol particle diffusivity Δ_{p} is computed as

$$\Delta_{\text{p}} = \frac{k_{\text{B}}Ts}{3\pi\nu D_{\text{g}}}$$ (B4)

with Boltzmann constant k_{B}, atmospheric dynamic viscosity $\nu = 1.8274 \times 10^{-5}\,\text{Pa}\,\text{s}\cdot(T/293.15\,\text{K})^{0.74}$, and the modal median diameter D_{g}. The factor

$$s = 1 + \left[2.514 + 0.8\exp\left(-0.55\frac{D_{\text{g}}}{\lambda_{\text{air}}}\right)\right]\frac{\lambda_{\text{air}}}{D_{\text{g}}},$$ (B5)

where λ_{air} stands for the mean free path of air, is known as the "slip correction". Combining the two scavenging rates (Eqs. B2-B3) one arrives at the total liquid nucleation scavenged fraction per unit time:

$$\eta_{\text{lnu}} = \eta_{\text{lnu}}^{\text{nuc}} + \eta_{\text{lnu}}^{\text{Br}} - \eta_{\text{lnu}}^{\text{nuc}}\,\eta_{\text{lnu}}^{\text{Br}}.$$ (B6)

The negative term accounts for the fact that aerosol particles cannot be scavenged at the same time by both, nucleation and Brownian motion.

Snow impaction scavenging is parameterized as

$$\eta_{\text{sim}} = \frac{\left[1 - \exp\left(-360\,\text{m}^2\,\text{s}\,\text{kg}^{-1}\right)\right]}{\Delta t},$$ (B7)

with the snow mass flux F_{s} per unit area and time. The rate coefficient η_{sim} is applied to particles of all MADE3 modes.

Finally, the rain impaction scavenging rate is assumed to be (Tost et al., 2006a)

$$\eta_{\text{rim}} = \frac{1}{\Delta t}\left[1 - \exp\left(-0.75\,F_{\text{r}}\,\Delta t \sum_{i=1}^{6}\frac{W_i E_i}{r_{\text{r},i}}\right)\right],$$ (B8)

where F_{r} is the rain mass flux per unit area and time. The terms $W_i E_i/r_{\text{r},i}$ are computed for six different values of the rain droplet radius $r_{\text{r},i}$ ($i = 1,\ldots,6$), namely for $0.1\,\text{mm}$, $0.2\,\text{mm}$, $0.5\,\text{mm}$, $1\,\text{mm}$, $2\,\text{mm}$, $5\,\text{mm}$. E_i is the collision efficiency of aerosol particles of size D_{g} and rain droplets of size $r_{\text{r},i}$ as parameterized by Slinn (1984), with weights W_i based on the rain droplet radii $r_{\text{r},i}$. Compared to measurements, E_i is likely underestimated for fine mode aerosol particles, which is a problem of any theoretically derived formulation for this parameter (Wang et al., 2010). As for snow impaction scavenging, the rain impaction scavenging rate is also applied to all MADE3 modes.

B2 Aerosol release from clouds and precipitation

In case of evaporation or sublimation of cloud particles or precipitating hydrometeors aerosol residues are released. The following section describes the algorithm for assigning the residual aerosol number and mass to the respective MADE3 modes. The mode naming convention for this section is described in the caption of Table 2. Corresponding numbers are as follows: ks = 1, km = 2, ki = 3, as = 4, am = 5, ai = 6, cs = 7, cm = 8, and ci = 9.

B2.1 Assignment of aerosol particle number concentrations

Let $\mathfrak{N}_q^{\mathrm{lnu}}$ and $\mathfrak{N}_q^{\mathrm{inu}}$ be the number concentrations of aerosol particles from mode q that were activated to form cloud droplets or ice crystals, respectively. Before all other calculations the insoluble ice nuclei number concentrations are assigned to the corresponding mixed modes (according to assumption 4 in Sect. 2.3) and the Aitken mode cloud particle cores to the corresponding

5   accumulation modes (according to assumption 5):

$$N_{\mathrm{as}}^{\mathrm{lnu}} = \mathfrak{N}_{\mathrm{as}}^{\mathrm{lnu}} + \mathfrak{N}_{\mathrm{ks}}^{\mathrm{lnu}}, \tag{B9a}$$

$$N_{\mathrm{as}}^{\mathrm{inu}} = \mathfrak{N}_{\mathrm{as}}^{\mathrm{inu}} + \mathfrak{N}_{\mathrm{ks}}^{\mathrm{inu}}, \tag{B9b}$$

$$N_{\mathrm{am}}^{\mathrm{lnu}} = \mathfrak{N}_{\mathrm{am}}^{\mathrm{lnu}} + \mathfrak{N}_{\mathrm{km}}^{\mathrm{lnu}}, \tag{B9c}$$

$$N_{\mathrm{am}}^{\mathrm{inu}} = \mathfrak{N}_{\mathrm{am}}^{\mathrm{inu}} + \mathfrak{N}_{\mathrm{ai}}^{\mathrm{inu}} + \mathfrak{N}_{\mathrm{km}}^{\mathrm{inu}} + \mathfrak{N}_{\mathrm{ki}}^{\mathrm{inu}}, \tag{B9d}$$

10   $$N_{\mathrm{cs}}^{\mathrm{lnu}} = \mathfrak{N}_{\mathrm{cs}}^{\mathrm{lnu}}, \tag{B9e}$$

$$N_{\mathrm{cs}}^{\mathrm{inu}} = \mathfrak{N}_{\mathrm{cs}}^{\mathrm{inu}}, \tag{B9f}$$

$$N_{\mathrm{cm}}^{\mathrm{lnu}} = \mathfrak{N}_{\mathrm{cm}}^{\mathrm{lnu}}, \tag{B9g}$$

$$N_{\mathrm{cm}}^{\mathrm{inu}} = \mathfrak{N}_{\mathrm{cm}}^{\mathrm{inu}} + \mathfrak{N}_{\mathrm{ci}}^{\mathrm{inu}}. \tag{B9h}$$

These operations are not fully compatible with the nucleation scavenging scheme of SCAV (Appendix B1) since the nucle-

15   ation scavenging rates considered by SCAV include not only nucleation of cloud droplets or ice crystals, but also Brownian motion scavenging within clouds. Separating the rates of the different processes for considering the pure nucleation rate in the assignment of cloud residues to the MADE3 modes would require fundamental and very extensive changes of the SCAV core algorithm, which would be far beyond the scope the present study. We therefore apply a simplified approach here: Brownian motion scavenging of interstitial aerosol in non-precipitating clouds is particularly important in case of ultrafine aerosols (e.g.,

20   Seinfeld and Pandis, 1998). Hence, large Brownian motion scavenging efficiencies of non-precipitating hydrometeors (nucleation scavenging operation in SCAV) can be expected in particular for the soluble Aitken mode, due to the small particle sizes of nucleating aerosol particles included in this mode. The insoluble and mixed Aitken modes as well as the accumulation and coarse modes contain larger particles, which are rather subject to nucleation scavenging than Brownian motion scavenging. Hence, we assume that only $\mathfrak{N}_{\mathrm{ks}}^{\mathrm{lnu}}$ and $\mathfrak{N}_{\mathrm{ks}}^{\mathrm{inu}}$ include major contributions of Brownian scavenging. This, therefore, might lead

25   to overestimations of the number of aerosol particles served as droplet or ice nuclei. In order to avoid this, we neglect $\mathfrak{N}_{\mathrm{ks}}^{\mathrm{lnu}}$ and $\mathfrak{N}_{\mathrm{ks}}^{\mathrm{inu}}$ in Eqs. (B9a) and (B9b), respectively. Due to its very small particle size, the soluble Aitken mode is only poorly activated to form cloud droplets and possible underestimations of $N_{\mathrm{as}}^{\mathrm{lnu}}$ due this simplification are probably small. In case of ice scavenging, 5% of N_{ks} is assumed to serve as ice nuclei in the present version of SCAV (Sect. 2.3 and Appendix B1). At high concentrations of ultrafine soluble particles this can lead to too high values of $N_{\mathrm{as}}^{\mathrm{inu}}$ compared to typical ice crystal

30   concentrations in pristine cirrus clouds. Neglecting $\mathfrak{N}_{\mathrm{ks}}^{\mathrm{inu}}$ has therefore an additional benefit under these conditions. In case of moderate number concentrations of the soluble Aitken mode, errors in the activated number $N_{\mathrm{as}}^{\mathrm{inu}}$ caused by neglecting $\mathfrak{N}_{\mathrm{ks}}^{\mathrm{inu}}$ are probably small compared to the uncertainties inherent in the assumption of constant 5% ice nucleation scavenging, since

variations of ice crystal number concentrations resulting from variing homogeneous freezing conditions (e.g., Kärcher and Lohmann, 2002) cannot be represented. Sensitivity simulations showed however that neglecting $\mathfrak{N}_{ks}^{lnu}$ and $\mathfrak{N}_{ks}^{inu}$ in Eqs. (B9a) and (B9b) has only a marginal effects on the results. This means that the effect of Brownian motion scavenging is small and further demonstrates that possible errors due to the simplified representations of nucleation scavenging are limited. Neverthe-

5 less, seperate budgeting of processes in the nucleation scavenging algorithm of SCAV could be the subject of future model development work, to enable a fully consistent representation of cloud residual aerosol.

The number concentrations denoted by N will be the output values of the mode assignment algorithm and are further modified as described in the following. According to Eqs. (B9a)-(B9h), cloud particle cores now belong to one of the four modes $q = 4, 5, 7, 8$ (corresponding to "as", "am", "cs", and "cm"). In the following, the symbol

10 $$N_q^{nuc} = N_q^{lnu} + N_q^{inu} \tag{B10}$$

is used to represent their cumulative number concentration in mode q. The total number concentration of cloud particles is assumed to remain constant, i.e., $\sum_{q=\{4,5,7,8\}} N_q^{nuc} = \text{const.}$ during the mode assignment process. In the present setup, SCAV later reduces all N_q^{nuc} to $10\,\%$ of their values to account for coagulation of cloud particles.

Now, the numbers of cloud particle cores per unit volume that are transferred to different aerosol modes upon evapora-
15 tion/sublimation of the cloud/precipitation particles are calculated. Such transfers are due to impaction scavenging of other aerosol particles. The symbol for the number of cores transferred from mode q to mode r per unit volume shall be N_q^r, the symbol for the total number of interstitial mode q particles per unit volume collected by impaction scavenging (Sect. B1) shall be $\mathfrak{N}_q^{imp}$.

Different cloud or precipitation particles can scavenge very different numbers of interstitial aerosol particles from the various
20 aerosol modes leading to a large variety of modifications of the original cloud particle cores. Representative mathematical descriptions of this complex system would lead to overly complicated and error-prone formulas to describe mode transfers of cores upon evaporation/sublimation of the cloud or precipitation particles. Hence, the system is simplified here by:

1. maximizing the transfer to mixed modes;

2. maximizing the transfer to coarse modes,

25 with a higher priority of the transfers to mixed modes. This deliberate "overestimation" of the transfer rates is motivated by the fact that the general reduction of N_q^{nuc} (to 10%, see above) does not account for transfers of cloud particle cores to other modes, since it simply reduces the total number of residues. The simplifications imply, for instance, the assumption that the interactions of interstitial aerosol with cloud particles are as homogeneously distributed as possible over the cloud particle population, maximizing the number of transfers induced by impaction scavenging.

30 It is acknowledged that these simplifications are somewhat arbitrary, but judging by previous simulation results – e.g., a relatively small long-term mean effect of cloud processing on aerosol particle aging when compared to condensation of trace gases – the associated error is expected to be tolerable. Nevertheless, different assumptions should be tested in the future in order to explore the sensitivity of the simulation results to different representations of the mode transfers.

No transfer is required out of mode "cm" as this mode represents the highest degree of aerosol particle mixing and aging. For modes "cs" and "am" the calculation is straightforward, as cores from these modes can only be transferred to mode "cm" (or remain in their respective mode). With the simplifying assumptions described above we obtain:

$$N_{\text{cs}}^{\text{cm}} = \text{MIN}\left(f_{\text{cs}}\left[\mathfrak{N}_{\text{km}}^{\text{imp}} + \mathfrak{N}_{\text{ki}}^{\text{imp}} + \mathfrak{N}_{\text{am}}^{\text{imp}} + \mathfrak{N}_{\text{ai}}^{\text{imp}} + \mathfrak{N}_{\text{cm}}^{\text{imp}} + \mathfrak{N}_{\text{ci}}^{\text{imp}}\right], \, N_{\text{cs}}^{\text{nuc}}\right), \tag{B11}$$

$$N_{\text{am}}^{\text{cm}} = \text{MIN}\left(f_{\text{am}}\left[\mathfrak{N}_{\text{cs}}^{\text{imp}} + \mathfrak{N}_{\text{cm}}^{\text{imp}} + \mathfrak{N}_{\text{ci}}^{\text{imp}}\right], \, N_{\text{am}}^{\text{nuc}}\right). \tag{B12}$$

Here,

$$f_q = \frac{N_q^{\text{nuc}}}{\sum_{r=\{4,5,7,8\}} N_r^{\text{nuc}}} \tag{B13}$$

is the fraction that mode q contributes to the total number concentration of cloud particle cores. The MIN operation is required because cloud or precipitation particles can collect multiple other particles via impaction scavenging, but their cores can of course only be transferred to mode "cm" once.

The situation is more complicated for the "as" cores, as they can be transferred to modes "am", "cs", and "cm", depending on the aerosol particles taken up by impaction scavenging. To simplify the system by maximizing the transfer to the mixed and/or coarse modes, the following assumptions are made:

– collected aerosol particles that contain insoluble material are distributed as evenly as possible over the "as" cores,

– fine, i.e., Aitken and accumulation mode particles that contain insoluble material are collected preferentially by cloud or precipitation particles that have not collected coarse mode particles which contain insoluble material,

– soluble coarse mode particles are collected preferentially by cloud or precipitation particles that have also collected fine mode particles that contain insoluble material.

Figure B1 may help the reader visualize these assumptions and the associated transfers described in the following.

Following the outlined scheme, the direct transfer of cloud or precipitation particle cores from mode "as" to mode "cm" by impaction of aerosol particles from modes "cm" and "ci" is considered first. Let

$$\Gamma = \text{MIN}\left(g_{\text{cm}} + g_{\text{ci}}, \, 1\right) \tag{B14}$$

be the fraction of mode r cores that has collected aerosol particles from the coarse modes that contain insoluble material (which is actually independent of r). Here,

$$g_q = \frac{f_r \mathfrak{N}_q^{\text{imp}}}{N_r^{\text{nuc}}} = \frac{\mathfrak{N}_q^{\text{imp}}}{\sum_{m=4,5,7,8} N_m^{\text{nuc}}} \tag{B15}$$

is the fraction of cores from mode r (or, in fact, from any mode) that has collected aerosol particles from mode q. The number of cores directly transferred from mode "as" to mode "cm" per unit volume can then be written as:

$$N_{\text{as}}^{\text{cm}} = \Gamma \, N_{\text{as}}^{\text{nuc}}. \tag{B16}$$
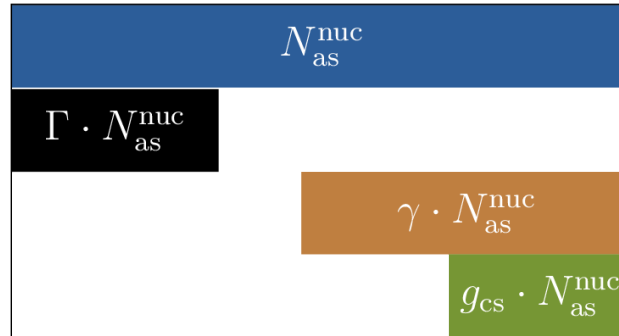
Figure B1. Illustration of the transfer concept for mode "as" cloud particle cores upon impaction scavenging of particles that induce such transfers. The blue bar represents the number concentration of mode "as" aerosol particles that have nucleated cloud particles, the boxes below it represent the number concentrations of impaction-scavenged aerosol particles from the three considered classes: coarse mode particles that contain insoluble material (black), fine mode particles that contain insoluble material (brown), and soluble coarse mode particles (green). See text for explanation of the symbols.

Subsequently, if any "as" cores remain, transfers from mode "as" to mode "cm" via collection of both fine mode particles that contain insoluble material and soluble coarse mode particles are considered. Let

$$\gamma = \mathrm{MIN}\big(g_{\mathrm{km}} + g_{\mathrm{ki}} + g_{\mathrm{am}} + g_{\mathrm{ai}},\, 1\big) \tag{B17}$$

be the analogue to Γ (Eq. B14) for collected fine mode particles that contain insoluble material. The expression for the transfer

5   of "as" cores to "cm" via collection of (at least) two aerosol particles (one from "km", "ki", "am", or "ai" and one from "cs") per cloud/precipitation particle then reads:

$$N_{\mathrm{as+2}}^{\mathrm{cm}} = \mathrm{MIN}\big(\gamma,\, g_{\mathrm{cs}},\, 1 - \Gamma\big)\, N_{\mathrm{as}}^{\mathrm{nuc}}. \tag{B18}$$

Once the terms for transfer to the mixed coarse mode have been established, and in case any "as" cores remain, transfer from "as" to "am" or "cs" without subsequent transfer to "cm" also has to be considered:

10   $$N_{\mathrm{as}}^{\mathrm{am}} = \mathrm{MAX}\Big(\mathrm{MIN}\big(\gamma,\, 1 - \Gamma\big) - g_{\mathrm{cs}},\, 0\Big)\, N_{\mathrm{as}}^{\mathrm{nuc}}, \tag{B19}$$

$$N_{\mathrm{as}}^{\mathrm{cs}} = \mathrm{MAX}\Big(\mathrm{MIN}\big(g_{\mathrm{cs}},\, 1 - \Gamma\big) - \gamma,\, 0\Big)\, N_{\mathrm{as}}^{\mathrm{nuc}}. \tag{B20}$$

After these calculations the N_q^{nuc} are redistributed accordingly (read the arrows as "new value on the left hand side is computed from old values on the right hand side"):

$$N_{\mathrm{as}}^{\mathrm{nuc}} \quad \rightarrow \quad N_{\mathrm{as}}^{\mathrm{nuc}} - N_{\mathrm{as}}^{\mathrm{am}} - N_{\mathrm{as}}^{\mathrm{cs}} - N_{\mathrm{as}}^{\mathrm{cm}} - N_{\mathrm{as+2}}^{\mathrm{cm}}, \tag{B21a}$$

15   $$N_{\mathrm{am}}^{\mathrm{nuc}} \quad \rightarrow \quad N_{\mathrm{am}}^{\mathrm{nuc}} + N_{\mathrm{as}}^{\mathrm{am}} - N_{\mathrm{am}}^{\mathrm{cm}}, \tag{B21b}$$

$$N_{\mathrm{cs}}^{\mathrm{nuc}} \quad \rightarrow \quad N_{\mathrm{cs}}^{\mathrm{nuc}} + N_{\mathrm{as}}^{\mathrm{cs}} - N_{\mathrm{cs}}^{\mathrm{cm}}, \tag{B21c}$$

$$N_{\mathrm{cm}}^{\mathrm{nuc}} \quad \rightarrow \quad N_{\mathrm{cm}}^{\mathrm{nuc}} + N_{\mathrm{as}}^{\mathrm{cm}} + N_{\mathrm{as+2}}^{\mathrm{cm}} + N_{\mathrm{am}}^{\mathrm{cm}} + N_{\mathrm{cs}}^{\mathrm{cm}}. \tag{B21d}$$

B2.2 Assignment of aerosol particle mass concentrations

As for the number concentrations, let $\mathfrak{m}_q^{\mathrm{lnu}}$ and $\mathfrak{m}_q^{\mathrm{inu}}$ be the mass concentrations of aerosol particles from mode q that were activated to form cloud droplets or ice crystals, respectively. Furthermore, let $\mathfrak{m}_q^{\mathrm{ch}}$ be the aerosol mass per unit volume generated within, or lost from, the cloud particles nucleated by mode q aerosol particles, which is due to cloud liquid-phase chemistry.

5 For consistency with the number treatment the insoluble ice nuclei mass concentrations are first assigned to the corresponding mixed modes and the Aitken mode cloud particle cores to the corresponding accumulation modes:

$$m_{\mathrm{as}}^{\mathrm{lnu}} = \mathfrak{m}_{\mathrm{as}}^{\mathrm{lnu}} + \mathfrak{m}_{\mathrm{ks}}^{\mathrm{lnu}}, \qquad\qquad m_{\mathrm{cs}}^{\mathrm{lnu}} = \mathfrak{m}_{\mathrm{cs}}^{\mathrm{lnu}}, \tag{B22a}$$

$$m_{\mathrm{as}}^{\mathrm{inu}} = \mathfrak{m}_{\mathrm{as}}^{\mathrm{inu}} + \mathfrak{m}_{\mathrm{ks}}^{\mathrm{inu}}, \qquad\qquad m_{\mathrm{cs}}^{\mathrm{inu}} = \mathfrak{m}_{\mathrm{cs}}^{\mathrm{inu}}, \tag{B22b}$$

$$m_{\mathrm{as}}^{\mathrm{ch}} = \mathfrak{m}_{\mathrm{as}}^{\mathrm{ch}} + \mathfrak{m}_{\mathrm{ks}}^{\mathrm{ch}}, \qquad\qquad m_{\mathrm{cs}}^{\mathrm{ch}} = \mathfrak{m}_{\mathrm{cs}}^{\mathrm{ch}}, \tag{B22c}$$

10 $$m_{\mathrm{am}}^{\mathrm{lnu}} = \mathfrak{m}_{\mathrm{am}}^{\mathrm{lnu}} + \mathfrak{m}_{\mathrm{km}}^{\mathrm{lnu}}, \qquad\qquad m_{\mathrm{cm}}^{\mathrm{lnu}} = \mathfrak{m}_{\mathrm{cm}}^{\mathrm{lnu}}, \tag{B22d}$$

$$m_{\mathrm{am}}^{\mathrm{inu}} = \mathfrak{m}_{\mathrm{am}}^{\mathrm{inu}} + \mathfrak{m}_{\mathrm{ai}}^{\mathrm{inu}} + \mathfrak{m}_{\mathrm{km}}^{\mathrm{inu}} + \mathfrak{m}_{\mathrm{ki}}^{\mathrm{inu}}, \qquad\qquad m_{\mathrm{cm}}^{\mathrm{inu}} = \mathfrak{m}_{\mathrm{cm}}^{\mathrm{inu}} + \mathfrak{m}_{\mathrm{ci}}^{\mathrm{inu}}, \tag{B22e}$$

$$m_{\mathrm{am}}^{\mathrm{ch}} = \mathfrak{m}_{\mathrm{am}}^{\mathrm{ch}} + \mathfrak{m}_{\mathrm{km}}^{\mathrm{ch}}, \qquad\qquad m_{\mathrm{cm}}^{\mathrm{ch}} = \mathfrak{m}_{\mathrm{cm}}^{\mathrm{ch}}. \tag{B22f}$$

Note that, in contrast to the treatment of particle numbers (Eqs. B9a and B9b), $m_{\mathrm{ks}}^{\mathrm{lnu}}$ and $m_{\mathrm{ks}}^{\mathrm{inu}}$ need to be considered here, since Brownian motion scavenging contributes to the mass of cloud particle residues while their number concentration does

15 not change. As before, the m will be the output values of our algorithm, the $\mathfrak{m}$ are the input values, computed by SCAV using the fractions given in Appendix B1. Note that there are no $\mathfrak{m}_{xi}^{\mathrm{ch}}$ ($x \in \{\mathrm{k, a, c}\}$) because ice phase chemistry is not considered.

In order to simplify the following expressions the cloud particle core mass concentration m_r^{nuc} for each mode $r = 4, 5, 7, 8$ ("as", "am", "cs", and "cm") is defined as the sum of the activated mode r aerosol particle mass per unit volume, the aerosol mass generated per unit volume within or on the cloud particles nucleated by mode r aerosol particles or lost from these cloud

20 particles, and the mass concentration of collected aerosol particles that do not induce transfers of the residual from mode r (let $\mathfrak{m}_q^{\mathrm{imp}}$ be the total mass concentration of mode q particles that are collected by impaction scavenging, Appendix B1):

$$m_{\mathrm{as}}^{\mathrm{nuc}} = m_{\mathrm{as}}^{\mathrm{lnu}} + m_{\mathrm{as}}^{\mathrm{inu}} + m_{\mathrm{as}}^{\mathrm{ch}} + f_{\mathrm{as}} \left[\mathfrak{m}_{\mathrm{ks}}^{\mathrm{imp}} + \mathfrak{m}_{\mathrm{as}}^{\mathrm{imp}} \right], \tag{B23a}$$

$$m_{\mathrm{am}}^{\mathrm{nuc}} = m_{\mathrm{am}}^{\mathrm{lnu}} + m_{\mathrm{am}}^{\mathrm{inu}} + m_{\mathrm{am}}^{\mathrm{ch}} + f_{\mathrm{am}} \left[\mathfrak{m}_{\mathrm{ks}}^{\mathrm{imp}} + \mathfrak{m}_{\mathrm{km}}^{\mathrm{imp}} + \mathfrak{m}_{\mathrm{ki}}^{\mathrm{imp}} + \mathfrak{m}_{\mathrm{as}}^{\mathrm{imp}} + \mathfrak{m}_{\mathrm{am}}^{\mathrm{imp}} + \mathfrak{m}_{\mathrm{ai}}^{\mathrm{imp}} \right], \tag{B23b}$$

$$m_{\mathrm{cs}}^{\mathrm{nuc}} = m_{\mathrm{cs}}^{\mathrm{lnu}} + m_{\mathrm{cs}}^{\mathrm{inu}} + m_{\mathrm{cs}}^{\mathrm{ch}} + f_{\mathrm{cs}} \left[\mathfrak{m}_{\mathrm{ks}}^{\mathrm{imp}} + \mathfrak{m}_{\mathrm{as}}^{\mathrm{imp}} + \mathfrak{m}_{\mathrm{cs}}^{\mathrm{imp}} \right], \tag{B23c}$$

25 $$m_{\mathrm{cm}}^{\mathrm{nuc}} = m_{\mathrm{cm}}^{\mathrm{lnu}} + m_{\mathrm{cm}}^{\mathrm{inu}} + m_{\mathrm{cm}}^{\mathrm{ch}} + f_{\mathrm{cm}} \sum_{q=1}^{9} \mathfrak{m}_q^{\mathrm{imp}}. \tag{B23d}$$

Now, the transferred mass concentrations are computed. For each mode $q = 4, 5, 7, 8$ the mass concentrations of particles that induced the transfer and a fraction of the core mass concentration m_q^{nuc} have to be transferred. As mode "cm" is only a target mode for residuals, there is no transfer out of this mode. All the mass that it receives by impaction scavenging stays in mode "cm". For the other three modes of cloud particle cores ("cs", "am", "as") the mass concentrations that have to be transferred

are calculated consistently with the number transfers:

$$m_{cs}^{cm} = f_{cs} \left[\mathfrak{m}_{km}^{imp} + \mathfrak{m}_{ki}^{imp} + \mathfrak{m}_{am}^{imp} + \mathfrak{m}_{ai}^{imp} + \mathfrak{m}_{cm}^{imp} + \mathfrak{m}_{ci}^{imp} \right] + \frac{N_{cs}^{cm}}{N_{cs}^{nuc}} m_{cs}^{nuc}, \tag{B24}$$

$$m_{am}^{cm} = f_{am} \left[\mathfrak{m}_{cs}^{imp} + \mathfrak{m}_{cm}^{imp} + \mathfrak{m}_{ci}^{imp} \right] + \frac{N_{am}^{cm}}{N_{am}^{nuc}} m_{am}^{nuc}, \tag{B25}$$

$$
\begin{aligned}
m_{as}^{cm} = & f_{as} \left[\mathfrak{m}_{cm}^{imp} + \mathfrak{m}_{ci}^{imp} \right] \\
& + \underbrace{\frac{\mathrm{MAX}\left(\Gamma + \gamma - 1,\, 0\right) N_{as}^{nuc}}{\gamma N_{as}^{nuc}}}_{\text{see note 1}} f_{as} \left[\mathfrak{m}_{km}^{imp} + \mathfrak{m}_{ki}^{imp} + \mathfrak{m}_{am}^{imp} + \mathfrak{m}_{ai}^{imp} \right] \\
& + \underbrace{\frac{\mathrm{MAX}\left(\Gamma + g_{cs} - 1,\, 0\right) N_{as}^{nuc}}{g_{cs} N_{as}^{nuc}}}_{\text{see note 2}} f_{as} \mathfrak{m}_{cs}^{imp} \\
& + \frac{N_{as}^{cm}}{N_{as}^{nuc}} m_{as}^{nuc},
\end{aligned}
\tag{B26}
$$

$$m_{as+2}^{cm} = \frac{N_{as+2}^{cm}}{\gamma N_{as}^{nuc}} f_{as} \left[\mathfrak{m}_{km}^{imp} + \mathfrak{m}_{ki}^{imp} + \mathfrak{m}_{am}^{imp} + \mathfrak{m}_{ai}^{imp} \right] + \frac{N_{as+2}^{cm}}{g_{cs} N_{as}^{nuc}} f_{as} \mathfrak{m}_{cs}^{imp} + \frac{N_{as+2}^{cm}}{N_{as}^{nuc}} m_{as}^{nuc}, \tag{B27}$$

$$m_{as}^{am} = \frac{N_{as}^{am}}{\gamma N_{as}^{nuc}} f_{as} \left[\mathfrak{m}_{km}^{imp} + \mathfrak{m}_{ki}^{imp} + \mathfrak{m}_{am}^{imp} + \mathfrak{m}_{ai}^{imp} \right] + \frac{N_{as}^{am}}{N_{as}^{nuc}} m_{as}^{nuc}, \tag{B28}$$

$$m_{as}^{cs} = \frac{N_{as}^{cs}}{g_{cs} N_{as}^{nuc}} f_{as} \mathfrak{m}_{cs}^{imp} + \frac{N_{as}^{cs}}{N_{as}^{nuc}} m_{as}^{nuc}. \tag{B29}$$

Notes:

1. fraction of collected fine mode particles containing insoluble material that is transferred with N_{as}^{cm},

2. fraction of collected soluble coarse mode particles that is transferred with N_{as}^{cm}.

Similar notes apply to the first terms in Eqs. (B27)-(B29). Note also that the terms with denominator γN_{as}^{nuc} add up to 1, and the terms with denominator $g_{cs} N_{as}^{nuc}$ do so as well. This is due to the limits imposed on the number transfers (see Sect. B2.1).

Finally, the mass concentration assignments are performed in an analogous manner as for the numbers (again, read the arrows as "new value on the left hand side is computed from old values on the right hand side"):

$$m_{\text{as}}^{\text{nuc}} \quad \rightarrow \quad m_{\text{as}}^{\text{nuc}} + f_{\text{as}} \left[\mathfrak{m}_{\text{km}}^{\text{imp}} + \mathfrak{m}_{\text{ki}}^{\text{imp}} + \mathfrak{m}_{\text{am}}^{\text{imp}} + \mathfrak{m}_{\text{ai}}^{\text{imp}} + \mathfrak{m}_{\text{cs}}^{\text{imp}} + \mathfrak{m}_{\text{cm}}^{\text{imp}} + \mathfrak{m}_{\text{ci}}^{\text{imp}} \right]$$
$$- m_{\text{as}}^{\text{am}} - m_{\text{as}}^{\text{cs}} - m_{\text{as}}^{\text{cm}} - m_{\text{as}+2}^{\text{cm}} \tag{B30a}$$
$$= m_{\text{as}}^{\text{nuc}} \left(1 - \frac{N_{\text{as}}^{\text{am}}}{N_{\text{as}}^{\text{nuc}}} - \frac{N_{\text{as}}^{\text{cs}}}{N_{\text{as}}^{\text{nuc}}} - \frac{N_{\text{as}}^{\text{cm}}}{N_{\text{as}}^{\text{nuc}}} - \frac{N_{\text{as}+2}^{\text{cm}}}{N_{\text{as}}^{\text{nuc}}} \right), \tag{B30b}$$

$$m_{\text{am}}^{\text{nuc}} \quad \rightarrow \quad m_{\text{am}}^{\text{nuc}} + f_{\text{am}} \left[\mathfrak{m}_{\text{cs}}^{\text{imp}} + \mathfrak{m}_{\text{cm}}^{\text{imp}} + \mathfrak{m}_{\text{ci}}^{\text{imp}} \right]$$
$$+ m_{\text{as}}^{\text{am}} - m_{\text{am}}^{\text{cm}} \tag{B30c}$$
$$= m_{\text{am}}^{\text{nuc}} \left(1 - \frac{N_{\text{am}}^{\text{cm}}}{N_{\text{am}}^{\text{nuc}}} \right) + m_{\text{as}}^{\text{am}}, \tag{B30d}$$

$$m_{\text{cs}}^{\text{nuc}} \quad \rightarrow \quad m_{\text{cs}}^{\text{nuc}} + f_{\text{cs}} \left[\mathfrak{m}_{\text{km}}^{\text{imp}} + \mathfrak{m}_{\text{ki}}^{\text{imp}} + \mathfrak{m}_{\text{am}}^{\text{imp}} + \mathfrak{m}_{\text{ai}}^{\text{imp}} + \mathfrak{m}_{\text{cm}}^{\text{imp}} + \mathfrak{m}_{\text{ci}}^{\text{imp}} \right]$$
$$+ m_{\text{as}}^{\text{cs}} - m_{\text{cs}}^{\text{cm}} \tag{B30e}$$
$$= m_{\text{cs}}^{\text{nuc}} \left(1 - \frac{N_{\text{cs}}^{\text{cm}}}{N_{\text{cs}}^{\text{nuc}}} \right) + m_{\text{as}}^{\text{cs}}, \tag{B30f}$$

$$m_{\text{cm}}^{\text{nuc}} \quad \rightarrow \quad m_{\text{cm}}^{\text{nuc}}$$
$$+ m_{\text{as}}^{\text{cm}} + m_{\text{as}+2}^{\text{cm}} + m_{\text{am}}^{\text{cm}} + m_{\text{cs}}^{\text{cm}}. \tag{B30g}$$

In the model, these expressions are applied to the individual species mass concentrations since they are the central prognostic quantities. This is legitimate since the individual species redistribute in proportion to the total mass.

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