Short Comments Received from GMD Executive Editor

We thank Lutz Gross for the comments. We have revised the manuscript following his comments. Our responses are compiled below (**SC:** short comment, **AR**: author's response):

SC: I would like to suggest the following changes to your manuscript for a revised submission:

- upload your program code and input files (if applicable) as a supplement
- clarify the license under which the reader can use the supplementary program code, see also <u>https://en.wikipedia.org/wiki/Creative Commons license</u> (this should be done in the section "Code availability" in your manuscript but it is also recommended to add to the header of any script)
- State in the section "Code availability" :
 - a persistent access to the code (preferable as supplement)
 - o license of use
 - the R versions required to run the scripts
 - the R version used to create the results shown in the paper

AR: We have included the codes in the supplement. We have reflected this change in the 'code availability' by including the license under which the codes can be used, the *R* version(s) that was used and is required to run the scripts.

The **code availability** section in the revised manuscript now reads:

"The *R* scripts used in this study are available in the supplemental material and can be found online at: <u>https://osf.io/a6q97/</u>. The *R* version 3.2.4 was used to create the scripts. A similar or higher version of *R* is required to run the scripts. The codes can be used, distributed, and reproduced for all non-commercial use under the terms of the Creative Commons Attribution-Noncommercial 3.0 Unported License (<u>http://creativecommons.org/licenses/by-nc/3.0/)</u>, provided the original work is properly cited."

Comments Received from Reviewer #1

We thank Ali Oskouie for carefully reviewing our manuscript and providing valuable suggestions. We have revised the manuscript following his comments. Our responses are compiled in the supplement, where we have addressed each comment individually and detailed the respective changes to the manuscript (*RC: reviewer's comment*, *AR: author's response*).

RC: The work presented here has special significance for researchers in terms of narrowing down their options based on parametrization made in this work. The authors have extensively discussed various variables involved in dry deposition process that looked like a textbook format. The authors should make every effort to shorten the text and eliminate some of the fundamental discussions as related to dry deposition along with the equations. There are some grammatical issues that needs to be addressed; however, it is minimal.

AR: We have omitted the first sentence of the introduction section and former Eq. (1), which was used to express the timescale for particle deposition. However, we think that it is necessary to keep the equations from five dry deposition parameterizations evaluated here in their original forms in the manuscript because they help guide the discussion of the relative complexity of the model parameterizations, which has been addressed in section 5 (lines 837-866) of the manuscript. We believe we have corrected the grammatical issues mentioned by the reviewer.

RC: The authors should avoid referencing to dry particle deposition, instead they should just mention 'dry deposition'.

AR: In the revised manuscript, after its first use in line 38 as 'dry particle deposition', we referred to it as 'dry deposition' in following texts when it appeared.

RC: Ambient particle density due to heterogeneity of particulate matter cannot be determined and used in these equations properly, so the authors should indicate such uncertainties.

AR: We understand your concern here. We addressed this issue by adding the following sentences (lines 394-397) in the revised manuscript.

"It is noted that a constant dry particle density of 1500 kg/m³ (Petroff and Zhang, 2010) was used in all Monte Carlo simulations. Because of the inhomogeneous nature of ambient particles, accurate quantification of particle density is challenging. In their work, Oskouie et al. (2003) developed methods using a time-of-flight instrument to minimize the effect of uncertainties in density estimation to particle size characterization."

RC: The authors should reference to Noll and Oskouie's pioneering work in the field of dry deposition to enrich their work with significant studies made in this field.

Comments Received from Reviewer #2

We thank Bruce Hicks for carefully reviewing our manuscript and providing valuable suggestions. We have revised the manuscript following his comments. Our responses are compiled in the supplement, where we have addressed each comment individually and detailed the respective changes to the manuscript (*RC: reviewer's comment*, *AR: author's response*).

RC: This is a well written paper, with conclusions that are important and with which I agree. The introduction gives one of the best overviews of the development of dry deposition formulations that I have come across. The text introduces a new (to me) method for ordering the sensitivity of a model output to the various input properties, and concludes (no surprise) that u* is a key property. Since u* is a measure of the level of turbulence affecting the eddy transport of all atmospheric properties, this is an expected result. However, it leads us further down a path that seems (to me) to have no foreseeable end, since u* is among the most difficult of all atmospheric properties to quantify, both instrumentally and even more so by using available models or even tabulations. The following observations arose during my reading of the mss.

AR: We understand and agree with your concern regarding difficulties in measuring friction velocity (u_*). We revised the opening paragraph of the conclusion section to reflect on your comment as follows (lines 878-883):

"One could argue that, if the different models exhibited greatest sensitivities to different parameters, and those parameters were more uncertain, the models exhibiting greatest sensitivity to the least certain parameters would be the most uncertain. In this way, sensitivity plays a potential role in determining which model is better. However, because our results showed that all models were most sensitive to u_* , or, at large size, *RH*, sensitivity does not end up playing a role in assigning which model is best. We also note that accurate measurement of u_* is extremely challenging (Andreas, 1992; Weber, 1999), and there exists ambiguity in its definition in boundary-layer meteorology (Weber, 1999)."

RC: The first sentence of the introduction could be omitted.

AR: We have omitted the first sentence.

RC: Line 33. Suggest omit "accurate

AR: We omitted the word as suggested (line 31) in the revised manuscript.

RC: Line 36. "particles" – plural.

AR: We have corrected the grammatical error (line 34) in the revised manuscript.

RC: Line 41. The time scale as stated is that which relates to gravitational settling. It is not appropriate for small particles, for which the relevant time scale needs to take into account the

depth of the layer in which the particles are mixed and the various resistances associated with their deposition. In any case, I do not see the need to refer to the time scale here.

AR: We agree with this comment. The equation (previously denoted as Eq. 1) stating the time scale for dry deposition is now omitted. The subsequent equations are re-numbered.

RC: Line 49. Terrain complexity is a central factor that none of the existing formulations Address.

AR: We agree that the issue of terrain complexity is not addressed in the five parameterizations we evaluated. To address this limitation, we have included the following sentences (lines 58-62):

"In all these models, the conventional resistance-based approach does not consider surface inhomogeneity or terrain complexity (i.e., deposition over flat terrain is assumed). However, Hicks (2008) argued about the importance of considering terrain complexity in dry deposition models because the assumption of surface homogeneity in existing deposition models limits the accuracy of pollutant load estimation in sensitive ecosystems that are located in complex terrain (e.g., on mountaintops or hills)."

RC: Line 57. " . . . there remain . . . " singular

AR: We have corrected the grammatical error (line 63) in the revised manuscript.

RC: Line 66. There is a recent paper in JGR (lead author – Hicks) that might shed some light on the way in which data have or have not agreed with model predictions.

AR: Thanks for pointing out to this reference. We have included the following sentences (lines 70-72):

"In a recent study, Hicks et al. (2016) compared five deposition models with measurements conducted over forests. They found that for particle sizes less than ca. 0.2 μ m, the modeled deposition velocities agreed fairly well with measured velocities. The largest discrepancy was observed for particle sizes of 0.3 to ca. 5.0 μ m."

RC: Lines 83-85. My immediate reaction is to wonder why the benefits of multi-variable partial correlation are not mentioned here. This attributes the variance among the contributing factors so that an ordering becomes obvious. The implication of the text is that a single correlation analysis is not appropriate because of the various covariances that could contribute. To my mind, this is precisely what a multiple regression is intended to resolve.

AR: For linear models, multiple variable partial correlation coefficients (PCC) provides the strength of linear relationship between the output and the parameter being tested for its sensitivity while removing the effects from all other input parameters (Brevault et al., 2013). Because the dry deposition models are non-linear, sensitivity analysis using PCC cannot provide accurate ranking of model input parameters in terms of their sensitivity to model outputs.

Rather, in non-linear models, a variance-based global sensitivity test approach such as Sobol' method for sensitivity analysis is more appropriate to achieve accurate parameter ranking and quantify the relative contributions of the input parameters to model outputs.

RC: Line 96 and onwards. This is the best description of what the models actually assume that I have come across. Congratulations to whoever it was who did the grunt work. Well done! The detail elevates my own fear that the modelers are constructing simulations that reflect articles of faith rather than of evidence. I appreciate that it is comparative examinations like that of the present text that might well shed some informed light on the benefits of model complexity. However, I have yet to see a reason not to start with something simple and add complexities as observations then warranted.

AR: Thanks for this comment.

RC: I recommend using readily apparent subheadings for each of the model descriptions.

AR: We have re-labeled the subheadings (sections 2.1 to 2.4) for each of the models in the revised manuscript.

RC: Section 3 – Methods. I note that the selection of data requires that each dataset contains measurements of all of the variables whose relevance is IMAGINED by the modelers. I consider this to be rather limiting. Surely, the selection of variables measured in field studies is a measure of what the experimentalists thought to be important factors. Examination of this might constitute a Delphic approach to determining what is important. I note, also, that requiring measurements of such things as leaf area index (LAI) assumes an application to situations in which a vegetated canopy of some specific kind dominates. For example, it has long been known that the hairs on leaves can play a role. How are these considered in the context of a LAI?

AR: Thank you for this comment. The presence of leaf hairs or trichromes is not considered in the context of LAI in any of the dry deposition models we evaluated in this study. We came across several measurement studies (Rauret et al. 1994; Howsam et al. 2000; Meyers et al. 2006; Burkhardt, 2010) that suggested enhanced particle deposition in the presence of trichomes. In fact, trichomes along with other microroughness elements such as epicuticular wax could increase fine particle deposition rates. We believe that the complexity associated with interactions between particle deposition and leaf microroughness elements poses a major challenge in dry deposition models to account for such interactions in a mechanistic fashion. That said, in our work, we stated that we did not consider the structural uncertainty (that which may arise from inadequate formulation of deposition processes) of the dry deposition models while evaluating their performance (lines 807-808).

RC: The definitions and roles of such properties as the zero plane displacement and the roughness length remain subjects of sometime heated discussion in meteorological circles. Using these concepts rather cavalierly in dry deposition models seems a poorly justified

extrapolation of a very poor basic understanding. I would welcome a more kindly-phrased reflection of this theme somewhere in the present text.

AR: To address your comment, we have included the following discussion (lines 796-807) in the revised manuscript:

"Additional uncertainties in the modeled deposition velocities could arise from inadequate model formulation and/or inappropriate use of certain micrometeorological parameters. For example, in dry deposition models (such as *PZ10*), *d* and z_0 are often calculated as a fraction of canopy height (*h*), and are often taken as $d \approx 2h/3$ and $z_0 \approx 0.1h$. These expressions are valid for dense canopies (Katul et al., 2010). If the leaf area density is highly skewed or shows a bimodal distribution, such approximations cannot be used (Katul et al., 2010). In addition, the parameter values of *d* and z_0 are subject to large uncertainty and are often very difficult to measure in urban areas (Cherin et al., 2015). Therefore, caution must be taken while using constant *d* and z_0 values from lookup tables. Also, current deposition models do not explicitly treat terrain complexity in their formulations. Hicks (2008) argued that conventional use of *d* and z_0 for non-flat terrain such as mountains is not appropriate in that context. In addition, experimentally derived values of *d* and z_0 often represent local characteristics. Thus, it poses a challenge to scale those up in a model grid cell (Schaudt and Dickinson, 2000) of atmospheric transport models. Using remote sensing, robust scaling of these parameter values is achieved, which could be used to acquire representative values in a model grid cell (Tian et al., 2011)."

RC: Line 345. The wind tunnel studies CANNOT be considered in the same context as open water deposition. One cannot get 2-m white-capped waves in a wind tunnel.

AR: We agree that the conditions are likely to be different between a wind tunnel and the field in the context of particle deposition over water surfaces. However, in dry deposition models (e.g., Petroff and Zhang, 2010, Zhang et al., 2001), the roughness length (z_0) for a water surface is calculated as a function of friction velocity and/or wind speed. The empirical expressions commonly used to estimate z_0 for water surfaces were developed based on laboratory or windwave tunnel experiments (Uneo and Deushi, 2003). For example, in the Petroff and Zhang, 2010 (PZ10) model, the z_0 was parameterized using an expression based on a wind-profile experiment conducted on a laboratory reservoir by Charnock (1955). In Charnock's expression, the roughness length does not depend on the wave state parameters such spectral peak frequency, wave steepness, and wave age (Uneo and Deushi, 2003). In our work, we used the formulations given by Keller and Keller (1992) to calculate roughness length for water surfaces, which were developed based on experiments conducted on a wave-tank. In their formulation, Keller and Keller (1992) also did not relate z_0 to wave state parameters (e.g., spectral peak frequency). They compared the wave-tank derived z_0 values with those obtained through open ocean momentum flux measurements conducted by Large and Pond (1981) and found good agreement. Given these reasons, we have included the wind-tunnel measurements in our study.

We changed the following sentence as "We note that the studies by Moller and Schumann, and Sehmel et al. were conducted in the wind tunnels, and thus the observed deposition does not (replaced *may not* with *does not*) necessarily reflect deposition under natural conditions" in the revised manuscript (line 361).

RC: Line 359. " . . . were used . . ."

AR: We have made the suggested correction (line 366) in the revised manuscript.

RC: Equations (63) and (64) are no more than expressions that quantify how much the averages differ. Why go into all of the mathematics when the description is so simple. And why not simply call it the amount by which the ratio of the averages differs from unity. (Defining BNMBF seems rather extreme. In practice, the distributions involved are likely to be log-normal or close to it, and hence the arithmetic average is an incorrect concept.)

AR: We have expressed these equations (now re-numbered as Eqs. 64-65) in their basic (simplified) forms. We have included a sentence to clarify the meaning of B_{NMFB} (lines 370-371) as suggested.

We agree that the distribution of deposition velocity (V_d) for a wide range of particle sizes (e.g., 0.01 -100 μ m) should exhibit a log-normal distribution (e.g., Figure 7 in Hicks et al., 2016). However, in most of the measurement studies used to evaluate model performance in our study, particle sizes (d_p) were confined to from 0.01 to 0.80 μ m, while the number of data points for d_p > 1.0 μ m (column 10 in Table 1) were minimal. Therefore, we think using B_{NMFB} is a reasonable approach to quantify the accuracy of the models.

RC: At this point, I encountered Tables 1 and 2. Some comments . . .

RC: What time of day are the numbers meant to represent?

AR: We are aware that the base values of certain parameters (e.g., relative humidity (*RH*), and Monin-Obukhov length (L_0)) could exhibit more diel variability than some other parameters listed in Table 2. To provide a basis for relative comparison, the base parameter values listed in Table 2 were used to perform uncertainty analysis, with the objective of demonstrating the overall uncertainties in modeled deposition velocities in the five parameterizations across five LUCs. These base values used for uncertainty analysis should be viewed as reasonable, arbitrary choices within the range of the reported values of those parameters in the measurement studies listed in Table 1. Here, we did not explicitly differentiate between daytime vs. nighttime conditions.

RC: What do the uncertainty ranges mean? If they are meant to refer to long-term ensemble averages base on some other variable, then I would agree with some of them.

AR: Each uncertainty range denotes the degree of imprecision (in percentage) associated with each of corresponding parameters listed in Table 2. To avoid confusion, we have referred to it

as 'uncertainty' in Table 2 in the revised manuscript. The uncertainty values (in percentage) were obtained from a literature search and the references are listed in column 3 of Table 2.

RC: The humidity assumption of 80% seems very high. Unless nighttime conditions are assumed.

AR: In several measurement studies (e.g., Neumann and den Hartog, 1985; Ruijgrok et al., 1997; Vong et al., 2004; Grönholm et al., 2009), particle deposition was measured at both daytime and nighttime. The threshold for very high humidity (*RH*) was reported as >95% (e.g., Ruijgrok et al., 1997) in the context of their measurement study. In addition, the ranges of *RH* were reported as 35-100% and 49-90% by Grönholm et al. (2009) and Vong et al. (2004), respectively. The choice of *RH* = 80% was arbitrary, but was within the reported range of *RH* by the measurement studies listed in Table 1. In addition, at high *RH* (e.g., *RH* close to 90%), atmospheric particles are likely to grow (the phenomena commonly known as hygroscopic growth) to twice their initial dry diameter (Vong et al., 2004). The hygroscopic growth was accounted for in the models according to Eq. (15).

RC: However, I note that nighttime is indeed assumed: the value of L is positive. A value of 50 m is commonly VERY stable.

AR: We understand your concern regarding the choice of Monin-Obukhov length (L_O) value as 50 m for the uncertainty test. In this study, we denoted uncertainties in the modeled V_d in normalized terms (i.e., the difference between the 95th and 5th percentiles V_d was divided by the 50th percentile V_d) rather than in absolute terms (i.e., the difference between the 95th and 5th percentiles V_d or the central 90% values of V_d). Therefore, the choice of L_O has minimal effect on the normalized uncertainties reported in Table 10.

RC: I could argue with all of the other "base values," all of which represent the environment that some investigator saw out his window. For example, grass 50 cm tall seems very unlikely. At that stage, it should have been harvested, grazed or mowed.

AR: The measurement studies used for model evaluation on grass reported typical values of grass height as 30-60 cm (e.g., Neumann and den Hartog, 1985; Vong et al., 2004). We agree that grass height of 50 cm could be higher than average height, but it was not uncommon in particle deposition measurement studies. For example, an article by Wesely et al. (1985) (not used in study) reported sulfur deposition on grass with height ranging from 40-50 cm.

RC: The use of a one-sided LAI worries me a little. Particles will deposit to both sides of leaves. Why use the one-sided value?

AR: We used one-sided LAI because atmospheric transport models (e.g., GEOS-Chem) regularly use MODIS derived LAI values as inputs to the dry deposition models. According to MODIS LAI database (<u>https://modis.gsfc.nasa.gov/data/dataprod/mod15.php</u>), "LAI is defined as one-sided green leaf area per unit ground area in broadleaf canopies and as half the total needle surface

area per unit ground area in coniferous canopies." In addition, reported values of LAI in the measurement studies (listed in Table 1) over vegetative surfaces often do not differentiate between two vs. one-sided LAI. Thus, to avoid ambiguity, we used one-sided LAI.

RC: In Table 2, I hope that the separate listing of ranges for u and u* does not mean that these are allowed to vary independently. The friction velocity is usually from 4% of wind speed (over water) to 20% of the wind speed (over a forest), with the variation of this proportionality being well known. It changes from day to night of course, but it seems that the analysis presented here is only for nighttime.

AR: The separate uncertainty ranges (now, for clarity, denoted as 'uncertainty' in Table 2, as compared to 'uncertainty range' in the previous manuscript) for wind speed (U) and friction velocity (u_*) denote respective imprecisions associated with those two parameters. In the uncertainty analysis presented in this paper, we used these two parameters independent of each other by allowing them to vary only within their specified imprecision/uncertainty ranges. The objective of the uncertainty analysis was to quantify the effects of imprecision in measured values of the input parameters (e.g., U, u_* , RH, L_0 , etc.) to the modeled V_d . Thus, we did not explicitly distinguish between daytime and nighttime periods.

RC: The variation of L from 10 to 100 m is very constraining, since once again the analysis represents conditions that would often be considered uncommonly stable.

AR: The analysis presented here assumed near-neutral conditions. The range (10 to 100 m) for L_0 (from stable side) was consistent with reported range for neutral conditions (-200 < L_0 < 200) by Pryor et al. (2007) (study #13 in Table 1). In addition, we did not find any change in parameter rankings reported in Table 11 by increasing the upper bound value of L_0 (in the assumed range of 10-100 m) by factors of 10 to 20.

RC: The roughness length quoted as the lower limit for snow/ice is 0.02 mm. As far as I know, this is less that aerodynamically smooth, and is therefore unlikely (some would say impossible). Please be careful.

AR: The lower limit of the roughness length (z_0) is consistent with the reported z_0 values in the measurement studies of particle deposition on ice/snow (e.g., Fig. 2a in Held et al., 2001a (study#24 in Table 1)). In addition, the article by Tjernström (2005) reported the lower bound z_0 value as 0.01 mm from Arctic Ocean Experiment of 2001.

RC: In Table 3, I note that all of the field programs are for daytime conditions, yet the Base Values tabulated indicate an assumption of stable (nighttime) conditions. Something must be wrong. A suitable explanation would be that the sign for *L* is incorrect.

AR: In several measurement studies listed in Table 1 (e.g., Neumann and den Hartog, 1985; Ruijgrok et al., 1997; Vong et al., 2004; Grönholm et al., 2009), particle deposition was measured under both daytime and nighttime conditions. The analysis presented here assumed stable conditions, thus the sign for *L*₀ is correct.

References

Brevault, L., Balesdent, M., Bérend, N., and Le Riche, R.: Comparison of different global sensitivity analysis methods for aerospace vehicle optimal design, 2013.

Burkhardt, J.: Hygroscopic particles on leaves: nutrients or desiccants?, Ecological Monographs, 80, 369-399, 2010.

Charnock, H.: Wind stress on a water surface, Quarterly Journal of the Royal Meteorological Society, 81, 639-640, 1955.

Cherin, N., Roustan, Y., Musson-Genon, L., and Seigneur, C.: Modelling atmospheric dry deposition in urban areas using an urban canopy approach, Geoscientific Model Development, 8, 893-910, 2015.

Grönholm, T., Launiainen, S., Ahlm, L., Mårtensson, E., Kulmala, M., Vesala, T., and Nilsson, E.: Aerosol particle dry deposition to canopy and forest floor measured by two-layer eddy covariance system, Journal of Geophysical Research: Atmospheres, 114, 2009.

Hicks, B. B.: On estimating dry deposition rates in complex terrain, Journal of Applied Meteorology and Climatology, 47, 1651-1658, 2008.

Hicks, B. B., Saylor, R. D., and Baker, B. D.: Dry deposition of particles to canopies—A look back and the road forward, Journal of Geophysical Research: Atmospheres, 121, 2016.

Howsam, M., Jones, K. C., and Ineson, P.: PAHs associated with the leaves of three deciduous tree species. I—concentrations and profiles, Environmental pollution, 108, 413-424, 2000.

Katul, G., Grönholm, T., Launiainen, S., and Vesala, T.: Predicting the dry deposition of aerosol-sized particles using layer-resolved canopy and pipe flow analogy models: Role of turbophoresis, Journal of Geophysical Research: Atmospheres, 115, 2010.

Keller, M. R., Keller, W. C., and Plant, W. J.: A wave tank study of the dependence of X band cross sections on wind speed and water temperature, Journal of Geophysical Research: Oceans, 97, 5771-5792, 1992.

Large, W. and Pond, S.: Open ocean momentum flux measurements in moderate to strong winds, Journal of physical oceanography, 11, 324-336, 1981.

Meyers, T., Luke, W., and Meisinger, J.: Fluxes of ammonia and sulfate over maize using relaxed eddy accumulation, Agricultural and forest meteorology, 136, 203-213, 2006.

Neumann, H. and Den Hartog, G.: Eddy correlation measurements of atmospheric fluxes of ozone, sulphur, and particulates during the Champaign intercomparison study, Journal of Geophysical Research: Atmospheres, 90, 2097-2110, 1985.

Petroff, A. and Zhang, L.: Development and validation of a size-resolved particle dry deposition scheme for application in aerosol transport models, Geoscientific Model Development, 3, 753-769, 2010.

Pryor, S., Larsen, S. E., Sørensen, L. L., Barthelmie, R. J., Grönholm, T., Kulmala, M., Launiainen, S., Rannik, Ü., and Vesala, T.: Particle fluxes over forests: Analyses of flux methods and functional dependencies, Journal of Geophysical Research: Atmospheres, 112, 2007.

Rauret, G., Llauradó, M., Tent, J., Rigol, A., Alegre, L., and Utrillas, M.: Deposition on holm oak leaf surfaces of accidentally released radionuclides, Science of the total environment, 157, 7-16, 1994.

Ruijgrok, W., Tieben, H., and Eisinga, P.: The dry deposition of particles to a forest canopy: a comparison of model and experimental results, Atmospheric Environment, 31, 399-415, 1997.

Schaudt, K. and Dickinson, R.: An approach to deriving roughness length and zero-plane displacement height from satellite data, prototyped with BOREAS data, Agricultural and Forest Meteorology, 104, 143-155, 2000.

Tian, X., Li, Z., Van der Tol, C., Su, Z., Li, X., He, Q., Bao, Y., Chen, E., and Li, L.: Estimating zero-plane displacement height and aerodynamic roughness length using synthesis of LiDAR and SPOT-5 data, Remote sensing of environment, 115, 2330-2341, 2011.

Tjernström, M.: The summer Arctic boundary layer during the Arctic Ocean Experiment 2001 (AOE-2001), Boundary-Layer Meteorology, 117, 5-36, 2005.

Ueno, K. and Deushi, M.: A new empirical formula for the aerodynamic roughness of water surface waves, Journal of oceanography, 59, 819-831, 2003.

Vong, R. J., Vickers, D., and Covert, D. S.: Eddy correlation measurements of aerosol deposition to short grass, Tellus B, 56, 105-117, 2004.

Weber, R. O.: Remarks on the definition and estimation of friction velocity, Boundary-Layer Meteorology, 93, 197-209, 1999.

Wesely, M., Cook, D., Hart, R., and Speer, R.: Measurements and parameterization of particulate sulfur dry deposition over grass, Journal of Geophysical Research: Atmospheres, 90, 2131-2143, 1985.

AR: Thank you for this reference. We have included the following sentence (lines 51-53) in the revised manuscript.

"Significant advances in understanding the governing mechanisms of dry deposition were made through use of experimental deposition data on walls of vertical pipes to develop size-resolved parameterizations for atmospheric particle deposition on ground surface (Muyshondt et al., 1998; Noll et al., 2001; Feng et al., 2008)."

RC: The uncertainties for ambient particulate density is addressed in Oskouie's work with unique calibration curves developed for determination of density of the ambient particles using supersonic TOF device which is used as the only calibration curves available for such characterization.

AR: Thanks for pointing out to the references.

References

Feng, J.: A size-resolved model and a four-mode parameterization of dry deposition of atmospheric aerosols, Journal of Geophysical Research: Atmospheres, 113, 2008.

Muyshondt, A., Anand, N., and McFarland, A. R.: Turbulent deposition of aerosol particles in large transport tubes, Aerosol Science and Technology, 24, 107-116, 1996.

Noll, K. E., Jackson, M. M., and Oskouie, A. K.: Development of an atmospheric particle dry deposition model, Aerosol Science & Technology, 35, 627-636, 2001.

Oskouie, A. K., Noll, K. E., and Wang, H.-C.: Minimizing the effect of density in determination of particle aerodynamic diameter using a time of flight instrument, Journal of aerosol science, 34, 501-506, 2003.

Petroff, A. and Zhang, L.: Development and validation of a size-resolved particle dry deposition scheme for application in aerosol transport models, Geoscientific Model Development, 3, 753-769, 2010.