Response to the topical editor (Slimane Bekki)

There are various tables providing boundary conditions for emissions. The authors should provide them under the form of ascii files with a link. That way, modellers who want to run the experiments just have to download the input files instead of having to recreate the files from the tables.

Dear Dr. Bekki,

Thank you very much for your comment. We have included now links to all input fields in the revised manuscript and our web page http://www.isamip.eu . This should ensure that all modellers could easily download the data in either ASCII or NETCDF format.

Response to reviewer 1

Thank you very much, for your very helpful comments and suggestions (indicated in bold and italic). You will find our point-by-point reply to them below.

"The authors discuss the climate impact of stratospheric volcanic aerosols, how their large scale distribution may be affected by stratospheric transport oscillations (QBO) and how their size distribution may change as a function of the injected SO2. A paragraph should be added, addressing the potential impact of the aerosol radiative interactions on some features of stratospheric dynamics and transport, as age of air and strat-trop exchange of trace species. Recent studies which may be relevant from this point of view, are those by Ray et al. (2014), Pitari et al. (2016a), Diallo et al. (2017). A brief paragraph on this aspect would make even stronger the need for the proposed MIP. This paragraph could probably be inserted in the Introduction or at the end of Subsection 3.3.2."

As the reviewer suggested we have included a couple of sentences on stratospheric aerosol and dynamics in the introduction and we also discuss uncertainties in mean age of air at the end of section 3.3.2.

Page 2, lines 57-65: "The consequent heating of the stratospheric aerosol layer strongly influences stratospheric dynamics amplifying the Brewer-Dobson circulation (BDC) and modifying the equator-to-pole temperature gradient. These two primary drivers cause changes to geostrophic zonal winds and the propagation of atmospheric waves (e.g. Bittner et al., 2016; Toohey et al., 2014) and lead to a strengthening of the polar vortex (e.g. Charlton-Perez et al., 2013). The heating from continued SO_2 injection to the stratosphere may further disturb or even "shut down" the quasi biennial oscillation (QBO) (e.g. Aquila et al., 2014; Niemeier and Schmidt, 2017). These composition-dynamics interactions also influence the transport and residence time of other long-lived species (N_2O, CH_4) (Pitari et al., 2016a; Visioni et al., 2017). The enhanced stratospheric aerosol layer after large volcanic eruptions causes also large mean age of air variations on time scales of several years (e.g. Ray et al., 2014; Muthers et al., 2016, Garfinkel et al., 2017)."

Page3, lines 78-81: "…..counteracted the warming due to increased greenhouse gases over that period (e.g. Solomon et al., 2011; Ridley et al., 2014; Santer et al., 2015). Small to moderate volcanic eruptions after 2008 also show an impact on the stratospheric circulation in the Northern Hemisphere, in particular on the pattern of decadal mean age variability and its trends during 2002–2011 (Diallo et al., 2017)."

Page 14, lines 487-497: "Analysing how the vertical profile of the enhanced stratospheric aerosol layer evolves during global dispersion and decay, will provide a key indicator for why the models differ, and what are the key driving mechanisms. Furthermore, the actual response of the BDC and mean age of air to Pinatubo is poorly constrained by existing reanalysis data (Garfinkel et al., 2017). While some modeling studies reported a decreasing mean age of air following volcanic eruptions throughout the stratosphere (Garcia et al., 2011; Garfinkel et al., 2017), show other studies an increase in mean age (Diallo et al., 2017). Moreover, Muthers et al. (2016) found decreasing age of air in the middle and upper stratosphere and increasing mean age below, while Pitari et al. (2016a) found decreasing mean age at higher levels of 30 hPa in the tropics and 10 hPa in the middle latitudes after the Pinatubo eruption. The HerSEA experiment in combination with a passive volcanic tracer might therefore help to better constrain the response of the BDC to volcanic eruptions using observations and help to clarify the uncertainties in age of air changes after the Pinatubo eruption. For all three major eruptions, we have identified key observational datasets (Table 7) that will provide benchmark tests to evaluate the vertical profile, covering a range of different aerosol metrics."

"References to new studies on volcanic aerosols may be added. The QBO impact on aerosol dispersal and e-folding time has been discussed in Pitari et al. (2016b) and could be cited at

page 5 line 181. A re-examination of the initial SO2 cloud lifetime was made in Mills et al. (2017) and could be cited at page 2 line 51*”*.

To take into account new developments/studies we have included a couple of recent published papers (indicated in blue) in the field:

Page 2, lines 59-52: “Major volcanic eruptions inject vast amounts of SO_2 into the stratosphere, which is converted into sulphuric acid aerosol with an e-folding time of about a month, which might be prolonged due to OH depletion within the dense SO_2 cloud in the first weeks following a large volcanic eruption (Mills et al., 2017)”.

Page 3, lines 90- 93: “For example, we now have a 2002-2012 long record of global altitude-resolved SO_2 ,and carbonyl sulphide (OCS) and aerosol volume density measurements provided by the Michelson Interferometer for Passive Atmospheric Sounding Environmental Satellite (MIPAS Envisat, Höpfner et al., 2013; 2015; Glatthor et al., 2015, Günther et al., 2018).”

Page 6, lines 193-196: “Internal variability associated with the QBO alters the isolation of the tropical stratosphere and subsequently the poleward transport of tropical stratospheric aerosol, thereby modulating its global dispersal, particle size distribution, and residence time (e.g. Trepte and Hitchmann, 1992; Hommel et al., 2015, Pitari et al., 2016b,).”

See also further new references included in the answers to other points by the reviewer.

“At page 7 lines 271-*273 the authors write: “Modelling groups are encouraged to include a* set of passive tracers to diagnose the atmospheric transport independently from emissions mostly following the CCMI recommendations (Eyring et al., 2013). These tracers are listed *in Table S3 in the supplementary material.” It should be specified that in case modelling* groups had already run these experiments, results produced and uploaded for CCMI may also be used for ISA-MIP, taking them directly from the CCMI data repository. I would also suggest to provide a link (as made in Eyring et al., 2013) where gridded input data may be available for download (S fluxes etc.).*”*

Thank you very much for your suggestion. This is a good point. We agree it will be valuable to compare the temporal variations in the ISA-MIP experiments (in particular the transient TAR and HErSEA experiments) with those from the CCMI REF-C1 and REF-C1SD simulations, which include the full fix of external forcing variations. Although we do not feel there is a need to specify this in the paper we will certainly encourage the leads for those two ISA-MIP experiments to consider this, and approach the relevant CCMI coordinators accordingly. In particular, there may still be time to double-check whether any extra diagnostics should be added to the ISA-MIP simulations as most groups will still be finalizing the final set-up for their integrations.

Forcings and other data sets will be made available on the ISA-MIP website: http://www.isamip.eu and through specific links, which will be included in the revised manuscript.

“At the beginning of Section 2 (page 5 lines 15-155) the following sentence sounds odd: *“However, the focus of the ISA*-MIP experiments described here is on comparing to measurements of the overall optical and physical properties of the stratospheric aerosol

***layer, which is manly determined by stratospheric aerosol". Maybe the final "aerosol"* should be substituted with sulfate.*"**

We have corrected the sentence to:

Page 5, lines 167-169: "However, the focus of the ISA-MIP experiments described here is on comparing to measurements of the overall optical and physical properties of the stratospheric aerosol layer, which is mainly determined by stratospheric aerosol sulphate."

"The discussion at the end of Section 2 (page 7 lines 210-219) could probably be made even more robust with reference to sulfate geoengineering studies. Some of these have *highlighted differences in what the authors themselves call "a crucial point", i.e.,* the different degree of isolation of the tropical pipe and the meridional transport of sulfate aerosols through the subtropical barrier. See for example Tilmes et al. (2015) and Visioni et al. (2018).*"*

We have added a sentence to refer to sulphate geoengineering studies

Page 7, lines 233-237: "Sulphate geoengineering studies confirm the importance of the model dependent meridional transport through the subtropical barrier (e.g. Niemeier and Timmreck, 2015; Visoni et al., 2018; Kleinschmitt et al., 2018). Reasons for these differences need to be understood with a multi-model comparison study, as suggested for example by Tilmes et al., (2015)."

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Response to reviewer 2

Thank you very much, for your very helpful comments and suggestions (indicated in bold and italic). You will find our point-by-point reply to them below.

"My only comment regarding the protocol concerns the TAR experiment "db1". **This is a mandatory (Tier 1) experiment that uses preferably tabulated point sources or, optionally, 3D data. However, the latter option is given an own identification in the protocol and a different priority (Tier 3). There may thus be a conflict, as a mandatory experiment is optionally bypassed by performing a low-priority experiment. If the two experiments are equivalent alternatives, they should be given the same priority, or even appear as the same experiment with the selected option to be reported in the metadata.**

I recommend some clarification."

We accept the original description of the TAR experiment may have been confusing. To ensure comparability between all of the three data sets we agreed on point sources for the volcanic emission in the mandatory (TIER1) experiments. For clarification we have changed the following sentence:

Page 11, lines 379-380: "If modelling groups prefer not to use point sources, we additionally offer VolcDB1_3D which provides a series of discrete 3D gridded SO_2 injections at specified times."

to

Page l1, lines 37-382: "To test the effect of the implementation strategy (point source vs cloud) an additional non-mandatory experiment has been set up: TAR_db1_sub with VolcDB1_3D as corresponding data set which provides a series of discrete 3D gridded SO_2 injections at specified times. "

To clarify, we also changed slightly the text:

Page 12, lines 401-402: "The optionally provided VolcDB1_3D data set, contains volume mixing ratio distributions of the injected SO_2 on a T42 Gaussian grid with 90 levels.

to

Page 12, lines 401-404: "The VolcDB1_3D data set, for the optional experiment tar_db1_3D contains volume mixing ratio distributions of the injected SO_2 cloud on a T42 Gaussian grid with 90 levels. The integral SO_2 mass for each injection is the same."

"As a general note, an expanded description about potential synergies and links with other ongoing MIPs would better highlight the value of ISA-MIP for the broader climate modelling community."

We have included in the summary the following sentences:

Pages 18-19, lines 632-649: "For example, the CMIP6 Geoengineering Model Intercomparison Project (GeoMIP, Kravitz et al., 2015) investigates common ways in which climate models treat various geoengineering scenarios some of them via sulphate aerosols (e.g. Tilmes et al., 2015). However, there is a large inter model spread for the cooling efficiency of sulphate aerosol, i.e. the normalized cooling rate per injected unit of sulphur (Moriyama et al., 2016). ISA-MIP is therefore of special importance for GeoMIP as it could help to understand the reason for these uncertainties, to better constrain the forcing efficiency and to improve future scenarios. Furthermore it is so far not clear whether the large inter-model spread of the CMIP5 models in the simulated post-volcanic climate response mostly depends on uncertainties in the imposed volcanic forcing or on an insufficient representation of climate processes. To discriminate the individual uncertainty factors it is useful to develop standardized experiments/model activities that systematically address specific uncertainty factors. Hence ISA-MIP, which covers the uncertainties in the pathway from the eruption source to the volcanic radiative forcing, will complement the CMIP6 VolMIP project (Zanchettin et al., 2016) which addresses the pathway from the forcing to the climate response and the feedback, by studying the uncertainties in the post-volcanic climate response to a well-defined volcanic forcing. ISA-MIP also complements the chemistry climate model initiative CCMI (Eyring et al., 2013) and the Aerosol Comparison (AeroCom) initiative (Schulz et al., 2006) as well as the Aerosol Chemistry Model Intercomparison Project (AerChemMIP, Collins et al., 2017) as it concentrates on stratospheric aerosol which is not in the focus of all these activities."

Specific comments on the manuscript:

Line 78: please check, the acronym OCS seems to be only introduced in line 164

We have included the explanation of OCS now earlier in the manuscript.

Page 3, lines 90-91: "we now have a 2002-2012 long record of global altitude-resolved SO_2, and carbonyl sulphide (OCS) and aerosol …"

Line 186: "compared to moderate eruptions"

We have revised the sentence accordingly to:

Page 6, lines 200-201: "… and predict a reduced cooling efficiency compared to moderate eruptions with moderate sulphur injections (e.g. Timmreck et al., 2010; English et al., 2013).

"Line 295: Across?"

We have revised the sentence to:

Page 9, line 313: "although new in situ measurements indicate that the cross-tropopause-SO2-flux is neglible over Mexico and central America (Rollins et al., 2017)."

"Line 321: the nudging period for the QBO is 1980-2000 (21 years) but the experiment only consists of 20 years. It seems that to include the year 2000 at the end of the simulation, the nudging period should start in 1981."

We have revised this accordingly:

Page 10, lines 329-340: .Modelling groups should run this simulation with varying QBO, either internally generated or nudged to the 1980-2000 period.

"Paragraph 3.3.3: It appears from Tables 5 and 6 that "VolcDB*" identify the datasets, whereas the experiment names are "TAR_db*/TAR_sub". It seems that the text in this **paragraph mixes the two (for instance in lines 374-376)."**

Thank you very much for this hint. We have revised the sentences to

Page 12, lines 394-397: "Summarising the number of experiments to be conducted within TAR: four are mandatory (TAR_base with no volcanic emission, Tar_db1/2/3), one additional is recommended (TAR-sub) and two others are optional (TAR_db4 and TAR_db1_3D; see Table 5 for an overview)."

Lines 432-438: this is certainly an interesting goal, but in this short description this appears at the edge or even slightly out of the scope of ISA-MIP itself. Can you expand on this?

Whilst we agree with the reviewer that there is no specific experiment aiming to understand the relationship between the ice core sulphate deposition and the stratospheric aerosol layer enhancements that drives the radiative cooling, the idea was to suggest that there is the potential for a systematic multi-model analysis of those 2 metrics (based on the HErSEA results) and seek to identify how uncertain historic volcanic forcings derived from ice core sulphate deposition may be.
We have added the following sentence to the revised manuscript stating that:

Page 13-14, lines, 440-463:“Although HErSEA has no specific experiment to understand the relationship between the ice core sulphate deposition and the stratospheric aerosol layer enhancements that drive the surface cooling, there is the potential for a systematic inter-model study (e.g. similar to Marshall et al., 2018) to identify how uncertain historic volcanic forcings derived from ice core sulphate deposition may be."

“Table 1: some of the information provided is not clearly described. For instance, are the numbers in parentheses in the “Total years” column the recommended integration years? This seems not to hold for the PoEMS where the numbers seem to refer to the number of perturbed parameters. The description of the number of specific experiments for PoEMS also seems to lack clarity.”

There was a mistake in the Table provided in the Discussions version of this article which we agree was confusing. In the revised manuscript we have changed Table 1 to be as shown below. We have also re-iterated (in the section 3.4.2 of the text, and in Table 1) the important requirement (currently only explained in the caption to Table 11) that the PoEMS parameter-scalings must only be applied in gridboxes with “volcanically-enhanced airmasses” (determined either by total-sulphur-vmr-threshold or the “passive Volc tracer”.

Page 17, lines 592-597: “When imposing the parameter-scalings, the models must only enact that change in grid boxes with volcanically-enhanced air masses. This can be determined either via total sulphur volume mixing ratio threshold suitable for the particular model, or via the “passive tracer Volc” recommended in section 3.3.3. Restricting the perturbation to the Pinatubo sulphur will leave pre-eruption conditions and tropospheric aerosol properties unchanged. ensuring a clean "uncertainty pdf" for the volcanic forcing."

Experiment	**Focus**	**Number of specific experiments**	**Years per experiment**	**Total years [A]**	**Knowledge-gap to be addressed**
….	….	……	…..	…..	………
Pinatubo Emulation in Multiple Models [PoEMS][B]	Perturbed parameter ensemble of runs to quantify uncertainty in each model’s predictions	10 experiments per parameter , where the number of parameters refers to the minimum (3), reduced (5) or standard (8) parameter set (see also Table 10)	3 per experiment[c]	90, 150 or 240	Intercompare Pinatubo perturbation to strat- aerosol properties with full uncertainty analysis over PPE run by each model. Quantify sensitivity of predicted Pinatubo perturbation stratospheric aerosol properties and radiative effects to uncertainties in injection settings and model processes Quantify and intercompare sources of uncertainty in simulated Pinatubo radiative forcing for the different complexity models.

[A] Each model will need to include an appropriate initialization and spin-up time for each ensemble member (~3-6 years depending on model configuration).

[B] As explained in the caption to Table 11 and section 3.4, models will need to restrict the PoEMS parameter-scaling to volcanically-enhanced air masses (either via total-sulphur-vmr threshold or passive volcanic SO_2 tracer)

[C] Although the Pinatubo enhancement to the stratospheric aerosol layer remained apparent until 1997 (e.g. Wilson et al., 2008), whereas the HErSEA experiments will continue longer, the PoEMS analysis will require only 3 post-eruption years to be run, as this gives sufficient time after the peak aerosol to characterize decay timescales robustly (e.g. ASAP2006, chapter5)

Table 1 General overview of the SSIRC ISA-MIP experiments

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List of the most relevant changes

- We have included links to all input fields.
- We have included a couple of sentences addressing the potential impact of the aerosol radiative interactions on some features of stratospheric dynamics and transport and we also discuss uncertainties in mean age of air now .
- We have included a couple of recent published papers.
- We have revised the description of the TAR experiment , TAR_db1 and TAR_db1_3D.
- We have revised the POEMS description of table 1.
- We have expanded the description about potential synergies and links with other ongoing MIPs.
- We have deleted Table S6 as it is included in the Brühl et al (2018) data reference.
- We have included a new table S6 which include a subset of volcanic emissions, that were derived based on the average mass of SO_2 emitted using VolcDB1, VolcDB2, and VolcDB3.

The Interactive Stratospheric Aerosol Model Intercomparison Project (ISA-MIP): Motivation and experimental design

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Abstract The Stratospheric Sulfur and its Role in Climate (SSiRC) interactive stratospheric aerosol model intercomparison project (ISA-MIP) explores uncertainties in the processes that connect volcanic emission of sulphur gas species and the radiative forcing associated with the resulting enhancement of the stratospheric aerosol layer. The central aim of ISA-MIP is to constrain and improve interactive stratospheric aerosol models and reduce uncertainties in the stratospheric aerosol forcing by comparing results of standardized model experiments with a range of observations. In this paper we present 4 co-ordinated inter-model experiments designed to investigate key processes which influence the formation and temporal development of stratospheric aerosol in different time periods of the observational record. The "Background" (BG) experiment will focus on microphysics and transport processes under volcanically quiescent conditions, when the stratospheric aerosol is controlled by the transport of aerosols and their precursors from the troposphere to the stratosphere. The "Transient Aerosol Record" (TAR) experiment will explore the role of small- to moderate-magnitude volcanic eruptions, anthropogenic sulphur emissions and transport processes over the period 1998-2012 and their role in the warming hiatus. Two further experiments will investigate the stratospheric sulphate aerosol evolution after major volcanic eruptions. The "Historical Eruptions SO_2 Emission Assessment" (HErSEA) experiment will

focus on the uncertainty in the initial emission of recent large-magnitude volcanic eruptions, while the "Pinatubo Emulation in Multiple models" (PoEMS) experiment will provide a comprehensive uncertainty analysis of the radiative forcing from the 1991 Mt. Pinatubo eruption.

1 Introduction

Stratospheric aerosol is an important component of the Earth system, which influences atmospheric radiative transfer, composition and dynamics, thereby modulating the climate. The effects of stratospheric aerosol on climate are especially evident when the opacity of the stratospheric aerosol layer is significantly increased after volcanic eruptions. Through changes in the radiative properties of the stratospheric aerosol layer, volcanic eruptions are a significant driver of climate variability (e.g. Myhre et al., 2013; Zanchettin et al., 2016). Major volcanic eruptions inject vast amounts of SO_2 into the stratosphere, which is converted into sulphuric acid aerosol with an e-folding time of about a month, which might be prolonged due to OH depletion within the dense SO_2 cloud in the first weeks following a large volcanic eruption (Mills et al., 2017).

. Observations show that the stratospheric aerosol layer remains enhanced for several years after major eruptions (SPARC, 2006). Such long-lasting volcanic perturbations cool the Earth's surface by scattering incoming solar radiation and warm the stratosphere by absorption of infrared solar and long-wave terrestrial radiation which affect the dynamical structure as well as the chemical composition of the atmosphere (e.g. Robock, 2000; Timmreck, 2012). The consequent heating of the stratospheric sulphate layer, impacts stratospheric dynamics in various ways. It amplifies the Brewer-Dobson circulation (BDC) and modifies the equator-to-pole temperature gradient, driving changes in geostrophic zonal winds and the propagation of atmospheric waves (e.g. Bittner et al., 2016; Toohey et al., 2014) and strengthening the polar vortex (e.g. Charlton-Perez et al., 2013). The heating from continued SO_2 injection to the stratosphere may further disturb or even "shut down" the quasi biennial oscillation (QBO) (e.g. Aquila et al., 2014; Niemeier and Schmidt, 2017). The radiatively driven changes also influence the transport and the lifetime of long-lived species (N_2O, CH_4) (Pitari et al., 2016a; Visioni et al., 2017). The enhanced stratospheric aerosol layer after large volcanic eruptions causes also large mean age of air variations on time scales of several years (e.g. Ray et al., 2014; Muthers et al., 2016, Garfinkel et al., 2017).

As the ocean has a much longer memory than the atmosphere, large volcanic eruptions could have a long lasting impact on the climate system that extends beyond the duration of the volcanic forcing (e.g., Zanchettin et al., 2012; Swingedouw et al., 2017). The chemical and radiative effects of the stratospheric aerosol are strongly influenced by its particle size distribution. Heterogeneous chemical reactions, which most notably lead to substantial ozone depletion (e.g. WMO Ozone Assessment 2007, chapter 3), take place on the surface of the stratospheric aerosol particles and are dependent on the aerosol surface area density. Aerosol particle size determines the scattering efficiency of the particles (e.g. Lacis et al., 1992).and their atmospheric lifetime (e.g., Pinto et al., 1989; Timmreck et al., 2010). Smaller-magnitude eruptions than 1991 Mt. Pinatubo eruption can also have significant impacts on climate. It is now established that a series of relatively small magnitude volcanic eruptions caused the increase in stratospheric aerosol observed between 2000 and 2010 over that period based on ground- and satellite-borne observations (Vernier et al., 2011b; Neely et al., 2013). Studies have suggested that this increase in stratospheric aerosol partly counteracted the warming due to increased greenhouse gases over that period (e.g. Solomon et al., 2011; Ridley et al., 2014; Santer et al., 2015). Small to moderate volcanic eruptions after 2008 also show an impact on the stratospheric circulation in the Northern Hemisphere, in particular on the pattern of decadal mean age variability and its trends during 2002–2011 (Diallo et al., 2017).

Since the 2006 SPARC Assessment of Stratospheric Aerosol Properties Report (SPARC 2006, herein referred as ASAP2006) the increase in observations of stratospheric aerosol and its precursor gases and in the number of models which treat stratospheric aerosol interactively, have advanced scientific understanding of the stratospheric aerosol layer and its effects on the climate (Kremser et al. 2016, herein referred to as KTH2016). In particular, research findings have given to the community a greater awareness of the role of the tropical tropopause layer (TTL) as a distinct pathway for transport into the stratosphere, of the interactions between stratospheric composition and dynamics, and of the importance of moderate-magnitude eruptions in influencing the stratospheric aerosol loading. In addition, over the last decade several new satellite instruments producing observations relevant to the stratospheric aerosol layer have become operational. For example, we now have a 2002-2012 long record of global altitude-resolved SO_2 ,~~and~~ carbonyl sulphide (OCS) and aerosol volume density measurements provided by the Michelson Interferometer for Passive Atmospheric Sounding Environmental Satellite (MIPAS Envisat, Höpfner et al., 2013; 2015; Glatthor et al., 2015, Günther et al., 2018). Furthermore aerosol extinction vertical profiles are available from limb-profiling instruments such as Scanning Imaging Absorption Spectrometer for Atmospheric Chartography (SCIAMACHY, 2002-2012; Bovensmann et al., 1999; von Savigny et al., 2015), Optical Spectrograph and InfraRed Imager System (OSIRIS, 2001-present, Bourassa et al., 2007), and Ozone Mapping and Profiler Suite–Limb Profiler (OMPS-LP, 2011-present, Rault and Loughman, 2013), and from the active sensor lidar measurements such as Cloud-Aerosol Transport System (CATS, 2015-present, Yorks et al., 2015) and Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP, 2006-present, Vernier et al., 2009). Existing measurements have become more robust, for example by homogenising the observations of aerosol properties derived from optical particle counter (OPC) and satellite measurements during stratospheric aerosol background periods (Kovilakam and Deshler, 2015), which previously showed large differences (Thomason et al., 2008). Other efforts include combining and comparing different satellite data sets (e.g. Rieger et al., 2015). However, some notable discrepancies still exist between different measurement datasets. For example, Reeves et al. (2008) showed that aircraft-borne Focused Cavity Aerosol Spectrometer (FCAS) measurements of the particle size distribution during the late 1990s yield surface area densities a factor 1.5 to 3 higher than that derived from Stratospheric Aerosol and Gases Experiment (SAGE-II) measurements.

On the modelling side there has been an increasing amount of global three-dimensional stratospheric aerosol models developed within the last years and used by research teams around the world (KTH2016). The majority of these global models explicitly simulate aerosol microphysical processes and treat the full life cycle of stratospheric aerosol, from the initial injection of sulphur containing gases, and their transformation into aerosol particles, to their final removal from the stratosphere. Several of these models also include the interactive coupling between aerosol microphysics, atmospheric chemistry, dynamics and radiation.

Given the improvements in observations and modelling of stratospheric aerosol since ASAP2006, we anticipate further advances in our understanding of stratospheric aerosol by combining the recent observational record with results from the current community of interactive stratospheric aerosol models. An Interactive Stratospheric Aerosol Model Intercomparison Project (ISA-MIP) has therefore been developed within the SSiRC framework. The SPARC activity Stratospheric Sulfur and its Role in Climate (SSiRC) (www.sparc-ssirc.org) was initiated with the goal of reducing uncertainties in the properties of stratospheric aerosol and assessing its climate forcing In particular, constraining simulations of historical eruptions with available observational datasets gives the potential to evaluate and substantially improve the accuracy of the volcanic forcing datasets used in climate models. This will not only enhance consistency with observed stratospheric aerosol properties and the underlying microphysical, chemical, and dynamical processes but also improve the conceptual understanding. The use of such new volcanic forcing datasets has the potential to increase the reliability of the simulated climate impacts of volcanic eruptions, which have been identified as a major influence on decadal global mean surface temperature trends in climate models (Marotzke and Forster, 2015).

The first international model inter-comparison of global stratospheric aerosol models was carried out within ASAP2006 and indicated that model simulations and satellite observations of stratospheric background aerosol extinction agree reasonably well in the visible wavelengths but not in the infrared. It also highlighted systematic differences between modelled and retrieved aerosol size, which are not able to detect the Aitken-mode sized particles (R<50nm) in the lower stratosphere (Thomason et al., 2008, Reeves et al., 2008; Hommel et al. 2011). While in ASAP2006, only five global two- and three-dimensional stratospheric aerosol models were included in the analysis, there are nowadays more than 15 global three-dimensional models worldwide available (KTH2016). No large comprehensive model intercomparison has ever been carried out to identify differences in stratospheric aerosol properties amongst these new interactive models. The models often show significant differences in terms of their simulated transport, chemistry, and removal of aerosols with inter-model differences in stratospheric circulation, radiative-dynamical interactions and exchange with the troposphere likely to play an important role (e.g. Aquila et al., 2012; Niemeier and Timmreck, 2015). The formulation of microphysical processes are also important (e.g. English et al. 2013), as are differing assumptions regarding the sources of stratospheric aerosols and their precursors. A combination of these effects likely explain the large inter-model differences as seen in Fig. 1 among global stratospheric aerosol models who participated in the Tambora intercomparison, a precursor to the "consensus volcanic forcings" aspects of the CMIP6 Model Intercomparison Project on the climatic response to Volcanic forcing (VolMIP, Zanchettin et al., 2016; Marshall et al., 2018~~7~~). Even for the relatively recent 1991 Mt. Pinatubo eruption, to reach the best agreement with observations, interactive stratospheric models have used a wide range of SO_2 injections amounts, from as low at 10 Tg of SO_2 (Dhomse et al., 2014; Mills et al., 2016) to as high as 20 Tg of SO_2 (e.g. Aquila et al., 2012; English et al., 2013).

Volcanic eruptions are commonly taken as a real-world analogue for hypothesised geoengineering via stratospheric sulphur solar radiation management (SS-SRM). Indeed many of the assumptions and uncertainties related to simulated volcanic perturbations to the stratospheric aerosol are also frequently given as caveats around research findings from modelling studies which seek to quantify the likely effects from SS-SRM (e.g. National Research Council, 2015), the mechanism-steps between sulphur injection and radiative cooling being common to both aspects (Robock et al., 2013). The analysis of the ISA-MIP experiments we expect to improve understanding of model sensitivities to key sources of uncertainty, to inform interpretation of coupled climate model simulations and the next Intergovernmental Panel on Climate Change (IPCC) assessment. It will also provide a foundation for co-operation to assess the atmospheric and climate changes when the next large-magnitude eruption takes place.

In this paper, we introduce the new ~~-~~model intercomparison project ISA-MIP developed within the SSiRC framework. In section 2 we provide an overview of the current state of stratospheric sulphur aerosol modelling and its greatest challenges. In section 3 we describe the scopes and protocols of the four model experiments planned within ISA-MIP. A concluding summary is provided in Section 4.

2. Modelling stratospheric aerosol; overview and challenges

Before we discuss the current state of stratospheric aerosol modelling and its greatest challenges in detail, we briefly describe the main features of the stratospheric sulphur cycle. We are aware of the fact that the stratospheric aerosol layer also contains organics and inclusions of meteoritic dust (Ebert et al., 2016) and, after volcanic events, also co-exists with volcanic ash (e.g. Pueschel et al., 1994: KTH2016). However, the focus of the ISA-MIP experiments described here is on comparing to measurements of the overall optical and physical properties of the stratospheric aerosol layer, which is mainly determined by ~~stratospheric aerosol~~ sulphate.

2.1 The stratospheric aerosol lifecycle

The stratospheric aerosol layer and its temporal and spatial variability are determined by the transport of aerosol and aerosol precursors in the stratosphere and their modification by chemical and microphysical processes (Hamill et al., 1997; ASAP2006; KTH2016). Volcanic eruptions can inject sulphur-bearing gases directly into the stratosphere which significantly enhances the stratospheric aerosol load for years. A number of observations show that stratospheric aerosol increased over the first decade of the 21st century (e.g. Hofmann et al., 2009; Vernier et al., 2011b; Ridley et al., 2014). Although such increase was attributed to the possible cause of Asian anthropogenic emission increase (Hofmann et al., 2009), later studies have shown that small-to-moderate magnitude volcanic eruptions are likely to be the major source of this recent increase (Vernier et al., 2011b; Neely et al., 2013; Brühl et al., 2015).

A stratospheric source besides major volcanic eruptions is the photochemical oxidation of ~~carbonyl sulphide (~~OCS~~)~~, an insoluble gas mainly inert in the troposphere. Tropospheric aerosols and aerosol precursor also enter the stratosphere through the tropical tropopause and through convective updrafts in the Asian and North

American Monsoons (Hofmann et al., 2009; Hommel et al., 2011; Vernier et al., 2011a; Bourassa et al., 2012; Yu et al., 2015). In the stratosphere, new sulphate aerosol particles are formed by binary homogenous nucleation (Vehkamäki et al., 2002), a process in which sulphuric acid vapour (H_2SO_4(g)) and water vapour condense simultaneously to form a liquid droplet. The condensation of H_2SO_4(g) onto pre-existing aerosol particles and the coagulation among particles shift the aerosol size distribution to greater radii. This takes place especially under volcanically perturbed conditions, when the concentrations of aerosol in the stratosphere are higher (e.g. Deshler et al., 2008).

From the tropics, where most of the tropospheric aerosol enters the stratosphere and the OCS chemistry is most active, the stratospheric aerosol particles are transported poleward within the large-scale ~~Brewer-Dobson circulation (~~BDC~~)~~ and removed through gravitational sedimentation and cross-tropopause transport in the extra-tropical regions. Internal variability associated with the ~~quasi biennial oscillation (~~QBO~~)~~ alters the isolation of the tropical stratosphere and subsequently the poleward ~~extra-tropical~~ transport of tropical~~he~~ stratospheric aerosol, and modifies its global dispersal ~~tribution~~, particle size distribution, and residence ~~life~~time (e.g. Trepte and Hitchmann, 1992; Hommel et al., 2015, Pitari et al., 2016b)~~.~~

In general, under volcanically perturbed conditions with larger amounts of injected SO_2, aerosol particles grow to much larger radii than in volcanic quiescent conditions (e.g. Deshler, 2008). Simulation of extremely large volcanic sulphur rich eruptions show a shift to particle sizes even larger than observed after the Pinatubo eruption, and predict a reduced cooling efficiency compared to moderate eruptions with moderate sulphur injections (e.g. Timmreck et al., 2010; English et al., 2013).

2.2 Global stratospheric aerosol models, current status and challenges

A comprehensive simulation of the spatio-temporal evolution of the particle size distribution is a continuing challenge for stratospheric aerosol models. Due to computational constraints, the formation of the stratospheric aerosol and the temporal evolution of its size distribution are usually parameterized with various degrees of complexity in global models. The simplest way to simulate the stratospheric aerosol distribution in global climate models is the mass only (bulk) approach (e.g. Timmreck et al., 1999a; 2003; Aquila et al., 2012), where only the total sulphate mass is prognostically simulated and chemical and radiative processes are calculated assuming a fixed typical particle size distribution. More complex methods are size-segregated approaches, such as the modal approach (e.g. Niemeier et al., 2009; Toohey et al., 2011; Brühl et al., 2012; Dhomse et al., 2014; Mills et al., 2016), where the aerosol size distribution is simulated using one or more modes, usually of log-normal shape. The mean radius of each mode of these size distributions varies in time and space. Another common approach is the sectional method (e.g. English et al., 2011; Hommel et al., 2011; Sheng et al., 2015a; for ref prior to 2006 see ASAP2006, chapter 5), where the particle size distribution is divided into distinct size sections. Number and width of the size sections are dependent on the specific model configuration, but are fixed throughout time and space. Size sections may be defined by an average radius, or by an average mass of sulphur, and are often spaced geometrically.

The choice of methods has an influence on simulated stratospheric aerosol size distributions and therefore on radiative and chemical effects. While previous model intercomparison studies in a box model (Kokkola et al., 2009) or in a two-dimensional framework (Weisenstein et al., 2007) were very useful for the microphysical schemes, they could not address uncertainties in the spatial transport pattern e.g. transport across the tropopause and the subtropical transport barrier, or regional/local differences in wet and dry removal. These uncertainties can only be addressed in a global three-dimensional model framework and with a careful validation with a variety of observational data.

The June 1991 eruption of Mt. Pinatubo, with the vast net of observations that tracked the evolution of the volcanic aerosol, provides a unique opportunity to test and validate global stratospheric aerosol models and their ability to simulate stratospheric transport processes. Previous model studies (e.g. Timmreck et al., 1999b; Aquila et al., 2012) highlighted the importance of an interactive online treatment of stratospheric aerosol radiative heating for the simulated transport of the volcanic cloud. A crucial point is the simulation of the tropical stratospheric aerosol reservoir (i.e., the tropical pipe, Plumb, 1996) and the meridional transport through the subtropical transport barrier. Some models show a very narrow tropical maximum in comparison to satellite data (e.g., Dhomse et al. 2014) while others show too fast transport to higher latitudes and fail to reproduce the long persistence of the tropical aerosol reservoir (e.g. Niemeier et al., 2009; English et al., 2013). Sulphate geoengineering studies affirm the importance of the model dependent meridional transport through the subtropical barrier (e.g. Niemeier and Timmreck, 2015; Visoni et al., 2018; Kleinschmidt et al., 2018). Reasons for these differences need to be understood with a multi-model comparison study, as suggested for example by Tilmes et al., (2015).

3. The ISA-MIP Experiments

Many uncertainties remain in the model representation of stratospheric aerosol. Figure 2 summarizes the main processes that determine the stratospheric sulphate aerosol mass load, size distribution and the associated optical properties. The four experiments in ISA-MIP are designed to address these key processes under a well-defined experiment protocol with prescribed boundary conditions (sea surface temperatures (SSTs), emissions). All simulations will be compared to observations to evaluate model performances and understand model strengths and weaknesses. The experiment "Background" (BG) focuses on microphysics and transport (section 3.1) under volcanically quiescent conditions, when stratospheric aerosol is only modulated by seasonal changes and interannual variability. The experiment "Transient Aerosol Record" (TAR) is addressing the role of time-varying SO_2 emission in particular the role of small- to moderate-magnitude volcanic eruptions and transport processes in the upper troposphere – lower stratosphere (UTLS) over the period 1998-2012 (section 3.2). Two further experiments investigate the stratospheric sulphate aerosol size distribution under the influence of large volcanic eruptions. "HErSEA" focuses on the uncertainty in the initial emission characteristics of recent large volcanic eruptions (section 3.3), while "PoEMS" provides an extensive uncertainty analysis of the radiative forcing of the Mt. Pinatubo eruption. In particular the ISA-MIP model experiments aim to address the following questions:

1. How large is the stratospheric sulphate load under volcanically quiescent conditions, and how sensitive is the simulation of this background aerosol layer to model specific microphysical parameterization and transport? (3.1)
2. Can we explain the sources and mechanisms behind the observed variability in stratospheric aerosol load since the year 2000? (3.2)
3. Can stratospheric aerosol observations constrain uncertainties in the initial sulphur injection amount and altitude distribution of the three largest volcanic eruptions of the last 100 years? (3.3)
4. What is the confidence interval for volcanic forcing of the Pinatubo eruption simulated by interactive stratospheric aerosol models and to which parameter uncertainties are the predictions most sensitive to? (3.4)

Table 1 gives an overview over all ISA-MIP experiments, which are described in detail below. In general each experiment will include several simulations from which only a subset is mandatory (Tier1). The modelling groups are free to choose in which of the experiments they would like to participate, however the BG Tier1 simulation is mandatory for all groups and the entry card for the ISA-MIP intercomparison. All model results will be saved in a consistent format (NETCDF) and made available via http://cera-www.dkrz.de/WDCC/ui, and compared to a set of benchmark observations. More detail technical information about data requests can be found in the supplementary material and on the ISA-MIP webpage: http://www.isamip.eu.

It is mandatory for participating models to run with interactive sulphur chemistry (see review in SPARC ASAP2006) in order to capture the oxidation pathway from precursors to aerosol particles, including aerosol growth due to condensation of H_2SO_4. Chemistry Climate Models (CCMs) with full interactive chemistry follow the Chemistry Climate Initiative (CCMI) hindcast scenario REF-C1 (Eyring et al. 2013, http://www.met.reading.ac.uk/ccmi/?page_id=11) for the treatment of chemical fields and emissions of greenhouse gases (GHGs), ozone depleting substances (ODSs), and very short-lived substances (VSLSs). Sea surface temperatures and sea ice extent are prescribed as monthly climatologies from the MetOffice Hadley Center Observational Dataset (Rayner et al. 2003). An overview of the boundary conditions is included in the supplementary material (Table S1). Table S2 reports the inventories to be used for tropospheric emissions of aerosols and aerosol precursors. Anthropogenic sulphur emissions and biomass burning are taken from the Monitoring Atmospheric Composition and Climate (MACC)-CITY climatology (Granier et al., 2011). S emissions from continuously erupting volcanoes are taken into account using Dentener et al. (2006) which is based on Andres and Kasgnoc (1998). OCS concentrations are fixed at the surface at a value of 510 pptv (Montzka et al., 2007; ASAP2006). If possible, DMS, dust, and sea salt emissions should be calculated online depending on the model meteorology. Models considering DMS oxidation should calculate seawater DMS emissions as a function of wind speed and DMS seawater concentrations. Otherwise, modelling groups should prescribe for these species their usual emission database for the year 2000. Each group can specify solar forcing for year 2000 conditions according to their usual dataset.

Modelling groups are encouraged to include a set of passive tracers to diagnose the atmospheric transport independently from emissions mostly following the CCMI recommendations (Eyring et al., 2013). These tracers are listed in Table S3 in the supplementary material. Models diagnose aerosol parameters as specified in Tables S4, S5. Additionally, volume mixing ratios of specified precursors are diagnosed

3.1 Stratospheric Background Aerosol (BG)

3.1.1. Summary of experiment

The overall objective of the BG experiment is to better understand the processes involved in maintaining the stratospheric background aerosol layer, i.e. stratospheric aerosol not resulting from direct volcanic injections into the stratosphere. The simulations prescribed for this experiment are time-slice simulations for the year 2000 with prescribed SST including all sources of aerosols and aerosol-precursors except for explosive volcanic eruptions. The result of BG will be a multi-model climatology of aerosol distribution, composition, and microphysical properties in absence of volcanic eruptions. By comparing models with different aerosol microphysics parameterization and simulations of background circulation with a variety of observational data (Table 2), we aim to assess how these processes impact the simulated aerosol characteristics.

3.1.2. Motivation

The total net sulphur mass flux from the troposphere into the stratosphere is estimated to be about 181 Gg S/yr based on simulations by Sheng et al. (2015a) using the SOCOL-AER model, 1.5 times larger than reported in ASAP2006 (KTH2016). This estimate, however, could be highly dependent on the specific characteristics of the model used, such as strength of convective systems, scavenging efficiency, and occurrence of stratosphere-troposphere exchange. Therefore, the simulated distribution of stratospheric background aerosol could show, especially in the lower stratosphere, a very large inter-model variability.

OCS is still considered the largest contributor to the aerosol loadings in the middle stratosphere. Several studies have shown that the transport to the stratosphere of tropospheric aerosol and aerosol precursors constitutes an important source of stratospheric aerosol (KTH2016 and references herein) although new in situ measurements indicate that the cross-tropopause-SO2-flux ~~the SO_2 flux cross the tropopause~~ is neglible over Mexico and central America (Rollins et al., 2017). Observations of the Asian Tropopause Aerosol Layer (ATAL, Vernier et al., 2011a) show that, particularly in the UTLS, aerosol of tropospheric origin can significantly enhance the burden of aerosol in the stratosphere. This tropospheric aerosol has a more complex composition than traditionally assumed for stratospheric aerosol: Yu et al. (2015), for instance, showed that carbonaceous aerosol makes up to 50% of the aerosol loadings within the ATAL. The rate of stratospheric-tropospheric exchange (STE) is influenced by the seasonality of the circulation and the frequency and strength of convective events in large-scale phenomena such as the Asian and North American monsoon or in small-scale phenomena such as strong storms. Model simulations by Hommel et al. (2015) also revealed significant QBO signatures in aerosol mixing ratio and size in the tropical middle stratosphere (Figure 3). Hence, the model specific implementation of the QBO (nudged or internally generated) could impact its effects on the stratospheric transport and, subsequently, on the stratospheric aerosol layer.

In this experiment, we aim to assess the inter-model variability of the background stratospheric aerosol layer, and of the sulphur mass flux from the troposphere to the stratosphere and vice versa. We will exclude changes in emissions and focus on the dependence of stratospheric aerosol concentrations and properties on stratospheric transport and stratosphere-troposphere exchange (STE). The goal of the BG experiment aims to understand how the model-specific transport characteristics (e.g. isolation of the tropical pipe, representation of the QBO and strength of convective systems) and aerosol parameterizations (e.g. aerosol microphysics and scavenging efficiency) affect the representation of the background aerosol.

3.1.3. Experiment setup and specifications

The BG experiment prescribes one mandatory (BG_QBO) and two recommended (BG_NQBO and BG_NAT) simulations (see Table 3). BG_QBO is a time slice simulation with conditions characteristic of the year 2000[1], with the goal of understanding sources, sinks, composition, and microphysical characteristics of stratospheric background aerosol under volcanically quiescent conditions. The time-slice simulation should be at least 20 year long, after a spin-up period of at least 10 years to equilibrate stratospheric relevant quantities such as OCS concentrations and age of air. The period seems to be sufficient to study differences in the aerosol properties but need to extended if dynamical changes e.g. in NH winter variability will be analysed. Modelling groups should run this simulation with varying QBO, either internally generated or nudged to the 1981~~0~~-2000 period.

If resources allow, each model should perform the sensitivity experiments BG_NQBO and BG_NAT. The specifics of these two experiments are the same as for BG_QBO, but BG_NQBO should be performed without varying QBO[2] and BG_NAT without anthropogenic emissions of aerosol and aerosol precursors, as indicated in Table S1. The goals of these sensitivity experiments are to understand the effect of the QBO on the background aerosol characteristics and the contribution of anthropogenic sources to the background aerosol loading in the stratosphere.

3.2 Transient Aerosol Record (TAR)

3.2.1 Summary of experiment

The aim of the TAR (Transient Aerosol Record) experiment is to investigate the relative contributions of volcanic and anthropogenic sources to the temporal evolution of the stratospheric aerosol layer between 1998 and 2012. Observations show that there is a transient increase in stratospheric aerosol loading, in particular after the year 2003, with small-to moderate-magnitude volcanic eruptions contributing significantly to this increase (e.g. Solomon et al., 2011, Vernier et al., 2011b; Neely et al., 2013; Ridley et al. 2014; Santer et al., 2015; Brühl et al., 2015). TAR model simulations will be performed using specified dynamics, prescribed sea surface

[1] To ensure comparability to the AeroCom simulations (http://aerocom.met.no/Welcome.html)

[2] Models with an internal generated QBO might nudge the tropical stratospheric winds.

temperature and time-varying SO_2 emissions. The simulations are suitable for any general circulation or chemistry transport models that simulate the stratospheric aerosol interactively and have the capability to nudge meteorological parameters to reanalysis data. The TAR protocol covers the period from January 1998 to December 2012, when only volcanic eruptions have affected the upper troposphere and lower stratosphere (UTLS) aerosol layer with SO_2 emissions about an order of magnitude smaller than Pinatubo. Time-varying surface emission datasets contain anthropogenic and natural sources of sulphur aerosol and their precursor species. The volcanic SO_2 emission inventories contain information of all known eruptions that emitted SO_2 into the UTLS during this period. It comprises the geolocation of each eruption, the amount of SO_2 emitted, and the height of the emissions. SO_2 emissions from continuously-degassing volcanoes are also included.

3.2.2 Experiment setup and specifications

Participating models are encouraged to perform up to seven experiments, based on five different volcanic SO_2 emission databases (hereafter referred to as VolcDB). Four experiments are mandatory, three other are optional. The volcanic experiments are compared to a reference simulation (~~noVolc~~ TAR_base) that does not use any of the volcanic emission databases, but emissions from continuously-degassing volcanoes. The aim of the reference simulation is to simulate the non-volcanically perturbed state of the stratospheric aerosol layer. In contrast to the experiment protocol BG (Section 3.1), here time-varying surface boundary conditions (SST/SIC) are applied, whereas BG intercompares model simulations under climatological mean conditions and uses constant 2000 conditions.

An overview of the volcanic emission inventories is given in Table 4 and in Figure 4 VolcDB1/2/3 are new compilations (Bingen et al., 201~~7~~; Neely and Schmidt, 2016; Carn et al., 2016), whereas a fourth inventory (VolcDB4; Diehl et al., 2012), provided earlier, for the AeroCom community modelling initiative, is optional. The databases use SO_2 observations from different sources and apply different techniques for the estimation of injection heights and the amount of emitted SO_2. The 4 inventories are provided in the form of tabulated point sources, with each modelling group to translate emitted SO_2 mass for each eruption into model levels spanning the upper and lower emission altitudes. To test the effect of the implementation strategy (point source vs cloud) an additional non-mandatory experiment has been set up: TAR_db1_3D with ~~If modelling groups prefer not to use point sources, we additionally offer~~ VolcDB1_3D~~-~~ as corresponding data set which provides a series of discrete 3D gridded SO_2 injections at specified times. In both versions of VolcDB1, the integral SO_2 mass of each injection is consistent.

We recommend performing one additional non-mandatory experiment TAR_sub in order to quantify and isolate the effects of 8 volcanic eruptions that either had a statistically significant effect on, for instance, tropospheric temperatures (Santer et al., 2014, 2015) or emitted significant amounts of SO_2 over the 1998 to 2012 time period. This experiment uses a subset of volcanic emissions (VolcDBSUB), that were derived based on the average mass of SO_2 emitted using VolcDB1, VolcDB2, and VolcDB3 for the following eruptions: 28 January 2005 Manam (4.0S, Papua New Guinea), 7 October 2006 Tavurvur (4.1 S, Papua New Guinea), 21 June 2009 Sarychev, (48.5° N, Kyrill, UDSSR) 8 November 2010 Merapi (7.3° S, Java, Indonesia), and 21 June 2011

Nabro (13.2° N, Eritrea). In addition the eruptions of Soufriere Hills (16.4° N, Monserrat) on 20 May 2006, Okmok (53.3° N, Alaska) on 12 July 2008 and Kasatochi (52.1° N, Alaska) on 7 August 2008 are considered (Table S6) although these are not discernible in climate proxy (Kravitz and Robock, 2010; Santer et al., 2014; 2015).

Summarising the number of experiments to be conducted within TAR: four are mandatory : ~~(~~TAR_base with no volcanic emission~~noVolc~~, TAR_db1/2/3,~~VolcDB1/2/3~~), one additional is recommended (TAR_sub ~~VolcDBSUB~~) and two others are optional (TAR_db4 ~~VolcDB4~~ and~~-~~ TAR_db1_3D~~VolcDB1_3D~~; see Table 5 for an overview).

Volcanic SO_2 Emission Databases

VolcDB1 (Bingen et al., 2017: Brühl (2018) Br~~-and Table S6~~) are updates from Brühl et al. (2015) using satellite data of MIPAS and OMI. For TAR, VolcDB1 has been extended based on data from Global Ozone Monitoring by Occultation of Stars (GOMOS), SAGE II, Total Ozone Mapping Spectrometer (TOMS), and the Smithsonian database. The ~~optionally provided~~ VolcDB1_3D data set,~~–~~ for the optional experiment TAR_db1_3D contains volume mixing ratio distributions of the injected SO_2 cloud on a T42 Gaussian grid with 90 levels. The integral SO_2 mass for each injection is the same. VolcDB2 (Mills et al., 2016; Neely and Schmidt, 2016) contains volcanic SO_2 emissions and plume altitudes for eruptions between that have been detected by satellite instruments including TOMS, OMI, OMPS, Infrared Atmospheric Sounding Interferometer (IASI), Global Ozone Monitoring Experiment (GOME/2), Atmospheric Infrared Sounder (AIRS), Microwave Limb Sounder (MLS) and the MIPAS instrument. The database is compiled based on published estimates of the eruption source parameters and reports from the Smithsonian Global Volcanism Program (http://volcano.si.edu/), NASA's Global Sulfur Dioxide Monitoring website (http://so2.gsfc.nasa.gov/) as well as the Support to Aviation Control Service (http://sacs.aeronomie.be/). The tabulated point source database also includes volcanic eruptions that emitted SO_2 into the troposphere only, as well as direct stratospheric emissions and has been used and compared to observations in Mills et al. (2016) and Solomon et al. (2016).

VolcDB3 uses the most recent compilation of the volcanic degassing data base of Carn et al. (2016). Observations from the satellite instruments TOMS, the High-resolution Infrared Sounder (HIRS/2), AIRS, OMI, MLS, IASI and OMPS are considered, measuring in the UV, IR and microwave spectral bands. Similar to VolcDB1/2, VolcDB3 also includes tropospheric eruptions.

Historically VolcDB4 is an older dataset, which relies on information from TOMS, OMI, the Global Volcanism Program (GVP), and other observations from the literature, covering the time period from 1979 to 2010. In contrast to the other inventories, VolcDB4 has previously been applied by a range of models within the AeroCom~~,~~ community (http://aerocom.met.no/emissions.html; Diehl et al., 2012, Dentener et al., 2006). Hence, it adds valuable information to the TAR experiments because it allows estimating how the advances in observational methods impact modelling results. It should be noted that VolcDB4 already contains the inventory of Andres and Kasgnoc (1998) for S emissions from continuously erupting volcanoes and should not be allocated twice when running this experiment.

Boundary Conditions, Chemistry and Forcings

To reduce uncertainties associated with model differences in the reproduction of synoptic and large-scale transport processes, models are strongly encouraged to perform TAR experiments with specified dynamics, where meteorological parameters are nudged to a reanalysis such as the ECMWF ERA-Interim (Dee et al., 2011). This allows models to reasonably reproduce the QBO and planetary wave structure in the stratosphere and to replicate as closely as possible the state of the BDC in the simulation period. Nudging also allows comparing directly to available observations of stratospheric aerosol properties (Table 2), such as the extinction profiles and AOD, and should enable the models to simulate the Asian tropopause layer (ATAL; Vernier et al., 2011a; Thomason and Vernier, 2013), which, so far, has been studied only by very few global models in great detail (e.g. Neely et al., 2014; Yu et al., 2015).

3.3. Historical Eruption SO2 Emission Assessment" (HErSEA)

3.3.1 Summary of experiment

This Historical Eruption SO_2 Emission Assessment (HErSEA) experiment will involve each participating model running a limited ensemble of simulations for each of the three largest volcanic perturbations to the stratosphere in the last 100 years: 1963 Mt. Agung, 1982 El Chichón and 1991 Mt. Pinatubo.

The main aim is to use a wide range of stratospheric aerosol observations to constrain uncertainties in the SO_2 emitted for each eruption (amount, injection height). Several different aerosol metrics will be intercompared to assess how effectively the emitted SO_2 translates into perturbations to stratospheric aerosol properties and simulated radiative forcings across interactive stratospheric aerosol CCMs with a range of different complexities. Whereas the TAR simulations (see section 3.2) use specified dynamics, and are suitable for chemistry transport models, for this experiment, simulations must be free-running with radiative coupling to the volcanically-enhanced stratospheric aerosol, thereby ensuring the composition-radiation-dynamics interactions associated with the injection are resolved. We are aware that this specification inherently excludes chemistry transport models, which must impose atmospheric dynamics. However, since the aim is to apply stratospheric aerosol observations in concert with the models to re-evaluate current best-estimates of the SO_2 input, and in light of the first order impact the stratospheric heating has on hemispheric dispersion from these major eruptions (e.g. Young, R. E. et al., 1994), we assert that this apparent exclusivity is entirely justified in this case.

As well as analysing and evaluating the individual model skill and identifying model consensus and disagreement for these three specific eruptions, we also seek to learn more about major eruptions which occurred before the era of satellite and in-situ stratospheric measurements. Our understanding of the effects from these earlier eruptions relies on deriving volcanic forcings from proxies such as sulphate deposition to ice sheets (Gao et al., 2007; Sigl et al., 2015; Toohey et al., 2013), from photometric measurements from astronomical observatories (Stothers, 1996, 2001) or from documentary evidence (Stothers, 2002; Stothers and Rampino, 1983; Toohey et al., 2016a). Although HErSEA has no specific experiment to understand the relationship between the ice core sulphate deposition and the stratospheric aerosol layer enhancements that drive the surface cooling, there is the potential for a systematic inter-model study (e.g. similar to Marshall et al., 2018) to identify how uncertain historic volcanic forcings derived from ice core sulphate deposition may be.

3.3.2 Motivation

In the days following the June 1991 Pinatubo eruption, satellite SO_2 measurements -show (e.g. Guo et al., 2004a) that the peak gas phase sulphur loading was 7 to 11.5 Tg [S] (or 14 -23 Tg SO_2). The chemical conversion to sulphuric aerosol that occurred in the tropical reservoir over the following weeks, and the subsequent transport to mid- and high-latitudes, caused a major enhancement to the stratospheric aerosol layer. The peak particle sulphur loading, through this global dispersion phase, reached only around half that in the initial SO_2 emission , the maximum particle sulphur loading measured as 3.7 to 6.7 Tg [S] (Lambert et al., 1993; Baran and Foot, 1994), based on an aqueous sulphuric acid composition range of 59 to 77% by weight (Grainger et al., 1993).

Whereas some model studies with aerosol microphysical processes find consistency with observations for SO_2 injection values of 8.5 Tg S (e.g., Niemeier et al., 2009; Toohey et al., 2011; Brühl et al., 2015), several recent microphysical model studies (Dhomse et al., 2014; Sheng et al. 2015a; Mills et al., 2016) find best agreement for an injected sulphur amount at, or even below, the lower end of the range from the satellite SO_2 measurements. Model predictions are known to be sensitive to differences in assumed injection height (e.g. Sheng et al., 2015b, Jones et al., 2016) and whether models resolve radiative heating and “self-lofting” effects also affects subsequent transport pathways (e.g. Young, R. E. et al., 1994; Timmreck et al. 1999b; Aquila et al., 2012). Another potential mechanism that could explain part of the apparent model-observation discrepancy is that a substantial proportion of the sulphur may have been removed from the plume in the first months after the eruption due to accommodation onto co-emitted ash/ice (Guo et al., 2004b) and subsequent sedimentation.

This ISA-MIP experiment will explore these issues further, with the participating models carrying out co-ordinated experiments of the three most recent major eruptions, with specified common SO_2 amounts and injection heights (Table 6). This design ensures the analysis can focus on key inter-model differences such as stratospheric circulation/dynamics, the impacts from radiative-dynamical interactions and the effects of aerosol microphysical schemes. -Analysing how the vertical profile of the enhanced stratospheric aerosol layer evolves during global dispersion and decay, will provide a key indicator for why the models differ, and what are the key driving mechanisms. Furthermore, the actual response of the BDC and mean age of air to Pinatubo is poorly constrained by existing reanalysis data (Garfinkel et al., 2017). While some modeling studies reported a decreasing mean age of air following volcanic eruptions throughout the stratosphere (Garcia et al., 2011; Garfinkel et al., 2017), show other an increase in mean age (Diallo et al., 2017). Moreover, Muthers et al. (2016) found decreasing mean age of air in the middle and upper stratosphere and increasing mean age below, while Pitari et al. (2016a) found decreasing mean age at higher levels of 30 hPa in the tropics and 10 hPa in the middle latitudes after the Pinatubo eruption. The HerSEA experiment in combination with a passive volcanic tracer might therefore help to better constrain the response of the BDC to volcanic eruptions using observations and help to clarify the uncertainties in age of air changes after the Pinatubo eruption.For all three major eruptions, we have identified key observational datasets (Table 7) that will provide benchmark tests to evaluate the vertical profile, covering a range of different aerosol metrics.

3.3.3 Experiment setup and specifications

Each modelling group will run a mini-ensemble of transient AMIP-type runs for the 3 eruptions with upper and lower bound SO_2 emissions and 3 different injection height settings: two shallow (e.g. 19-21 km and 23-25 km) and one deep (e.g. 19-25 km) (see Table 7). The seasonal cycle of the ~~Brewer Dobson circulation~~ BDC affects the hemispheric dispersion of the aerosol plume (e.g. Toohey et al., 2011) and the phase of the QBO is also known to be key control for tropical eruptions (e.g. Trepte and Hitchman, 1992). In order t~~T~~o quantify the contribution of the tracer transport, a passive tracer Volc (Table S3) will is recommended to be additionally initialized and transported. Note since the AMIP-type simulations will be transient, prescribing time-varying sea-surface temperatures, the models will automatically match the surface climate state (ENSO, NAO) through each post-eruption period. Where possible, models should re-initialise (if they have internally generated QBO) or use specified dynamics approaches (e.g. Telford et al., 2008) to ensure the model dynamics is consistent with the QBO evolution through the post-eruption period. General circulation models should use GHG concentrations appropriate for the period and models with interactive stratospheric chemistry should ensure the loading of Ozone Depleting Substances (ODSs) matches that for the time period.

Table 8 shows the settings for the SO_2 injection for each eruption~~.~~ Note that experience of running interactive stratospheric aerosol simulations shows that the vertical extent of the enhanced stratospheric aerosol will be different from the altitude range in which the SO_2 is injected. So, these sensitivity simulations will allow to assess the behaviour of the individual models with identical settings for the SO_2 injection.

For these major eruptions, where the perturbation is much larger than in TAR, model diagnostics include AOD and extinction at multiple wavelengths and heating rates (K/day) in the lower stratosphere to identify the stratospheric warming induced by simulated volcanic enhancement, including exploring compensating effects from other constituents (e.g. Kinne et al., 1992). To allow the global variation in size distribution to be intercompared, models will also provide 3D-monthly effective radius, with also cumulative number concentration at several size-cuts for direct comparison to balloon measurements. Examining the co-variation of the particle size distribution with variations in extinction at different wavelengths will be of particular interest in relation to approaches used to interpret astronomical measurements of eruptions in the pre-in-situ era (Stothers, 1996, 2001). A 3-member ensemble will be submitted for each different injection setting.

3.4. Pinatubo Emulation in Multiple models" (PoEMs)

3.4.1 Summary of experiment

The PoEMS experiment will involve each interactive stratospheric aerosol model running a perturbed parameter -ensemble (PPE) of simulations through the 1991-1995 Pinatubo-perturbed period. Variation-based sensitivity analysis will derive a probability distribution function (PDF) for each model's predicted Pinatubo forcing, following techniques applied successfully to quantify and attribute sources of uncertainty in tropospheric aerosol forcings (e.g. Carslaw et al., 2013). The approach will teach us which aspects of the radiative forcing from major eruptions is most uncertain, and will enable us to identify how sensitive model predictions of key features (e.g. timing and value of peak forcing and decay timescales) are to uncertainties in several model parameters. By comparing the time-signatures of different underlying aerosol metrics (mid-visible AOD, effective radius, particle number) between models, and crucially also against observations, may also help to reduce the natural forcing uncertainty, potentially thereby making the next generation of climate models more robust.

3.4.2 Motivation

The sudden global cooling from major eruptions is a key signature in the historical climate record and a natural global warming signature occurs after peak cooling as volcanic aerosol is slowly removed from the stratosphere. Quantitative information on the uncertainty range of volcanic forcings is therefore urgently needed. The amount of data collected by satellite-, ground-, and air-borne instruments in the period following the 1991 eruption of Mount Pinatubo (see e.g. section 3.3.2, Table 7) provides an opportunity to test model capabilities in simulating large perturbations of stratospheric aerosol and their effect on the climate. Recent advances in quantify uncertainty in climate models (e.g. Rougier et al., 2009;Lee at al. 2011) involve running ensembles of simulations to systematically explore combinations of different external forcings to scope the range of possible realisations. There are now a large number of general circulation models (GCMs) with prognostic aerosol modules, which tend to assess the stratospheric aerosol perturbation through the Pinatubo-perturbed period (see Table 9). Although these different models achieve reasonable agreement with the observations, this consistency of skill is achieved with considerable diversity in the values assumed for the initial magnitude and distribution of the SO_2 injection. The SO_2 injections prescribed by different models range from 5Tg-S to 10 Tg-S, and the upper edge of the injection altitude varies among models from as low as 18km to as high as 29km, as shown in Table 9. Such simulations also differ in the choice of the vertical distribution of SO_2 injection (e.g. uniform, Gaussian, or triangular distributions) and the horizontal injection area (one to several grid boxes). The fact that different choices of injection parameters lead to similar results in different models points to differences in the models' internal treatment of aerosol evolution. Accurately capturing microphysical processes such as coagulational, growth and subsequent rates of sedimentation has been shown to be important for volcanic forcings (English et al., 2013), but some studies (e.g. Mann et al., 2015) identify that these processes interplay also with aerosol-radiation interactions, the associated dynamical effects changing the fate of the volcanic sulphur and its removal into the troposphere. The PoEMS experiment will specifically assess this issue by adjusting the rate of specific microphysical processes in each model simultaneously with perturbations to SO_2 emission and injection height, thereby assessing the footprint of their influence on subsequent volcanic forcing in different complexity aerosol schemes and the relative contribution to uncertainty from emissions and microphysics.

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3.4.3 Experiment setup and specifications

For each model, an ensemble of simulations will be performed varying SO_2 injection parameters and a selection of internal model parameters within a realistic uncertainty distribution. A maximin Latin hypercube sampling strategy will be used to define parameter values to be set in each PPE member in order to obtain good coverage of the parameter space. The maximin Latin hypercube is designed such that the range of every single parameter is well sampled and the sampling points are well spread through the multi-dimensional uncertainty space – this is achieved by splitting the range of every parameter into N intervals and ensuring that precisely one point is in each interval in all dimensions, where N is the total number of model simulations, and the minimum distance between any pair of points in all dimensions is maximised. Fig. 6 shows the projection onto two dimensions of a Latin hypercube built in 8 dimensions with 50 model simulations. The size of the Latin hypercube needed will depend on the number of model parameters to be perturbed; the number of simulations to be performed will be equal to ~~seven~~ ten times the number of parameters - seven per parameter to build the emulator and three per parameter to validate the emulator. All parameters are perturbed simultaneously in the Latin hypercube.

In order to be inclusive of modelling groups with less computing time available, and different types of aerosol schemes, we define 3 options of experimental design with different numbers of perturbed parameters and thus simulation ensemble members. The 3 options involve varying all 8 (standard set), 5 (reduced set), or 3 (minimum set) of the list of uncertain parameters, resulting in ensembles of 80 ~~64~~ (standard), 5~~4~~0 (reduced) or 30 ~~24~~ (minimum) PPE members. The parameters to be varied are shown in Table 10, and include variables related to the volcanic injection, such as its magnitude, height, latitudinal extent, and composition, and to the life cycle of the volcanic sulphate, such as the sedimentation rate, its microphysical evolution, and the SO_2 to SO_4^{2-} conversion rate.

Prior to performing the full PPE, modelling groups are encouraged to run "One-At-a-Time" (OAT) test runs with each of the process parameters increased/decreased to its maximum/minimum value. Submission of these OAT test runs is encouraged (following the naming convention in Table 11) because as well as being an important check that the model parameter-scaling is being implemented as intended, the results will also enable intercomparison of single-parameter effects between participating models ahead of the full ensemble. When imposing the parameter-scalings, the models must only enact that change in gridboxes with volcanically-enhanced air masses. This can be determined either via total sulphur volume mixing ratio threshold suitable for the particular model, or via the "passive tracer Volc" recommended in section 3~~x~~.3~~x~~.3.~~–~~ Restricting the perturbation to the Pinatubo sulphur will leave pre-eruption conditions and tropospheric aerosol properties unchanged. ensuring a clean "uncertainty pdf" for the volcanic forcing"

~~-~~That this restriction to the parameter-scalings is operational is an important preparatory exercise and will need to have been verified when running the OAT test runs.

Once a modelling group has performed the PPE of simulations as defined by the Latin hypercube a statistical analysis will be performed. Emulators for each of a selection of key metrics will be built, following the approach described by Lee et al. (2011), to examine how the parameters lead to uncertainty in key features of the Pinatubo-perturbed stratospheric aerosol. The emulator builds a statistical model between the ensemble design and the key model output and once validated allows sampling of the whole parameter space to derive a PDF of each key model output.

Variance-based sensitivity analysis will then be used to decompose the resulting probability distribution into its sources providing information on the key sources of uncertainty in any model output. The two sensitivity indices of interest are called the main effect and the total effect. The main effect measures the percentage of uncertainty in the simulated metric due to each parameter-variation individually. The total effect measures the percentage of uncertainty in the key model output due to each parameter, including the additional contribution from its interaction with other uncertain parameters. The sources of model parametric uncertainty (i.e. the sensitivity indices) will be identified for each model with discussion with each group to check the results. By then comparing the sensitivity to the uncertain parameters across the range of participating models, we will learn about how the model's differing treatment of aerosol processes, and the inherent dynamical and chemical processes resolved in the host model, together determine the uncertainty in its predicted Pinatubo radiative forcings.

The probability distribution of observable key model outputs will also be compared to observations, in order to constrain the key sources of uncertainty and thereby reduce the parametric uncertainty in individual models. The resulting model constraints will be compared between models providing quantification of both parametric uncertainty and structural uncertainty for key variables such as AOD, effective radius and radiative flux anomalies. This sensitivity analysis will also identify the variables for which better observational constraints would yield the greatest reduction in model uncertainties.

4. Conclusions

The ISA-MIP experiments will improve understanding of stratospheric aerosol processes, chemistry, and dynamics, and constrain climate impacts of background aerosol "variability", small volcanic eruptions, and large volcanic eruptions. The experiments will also help to resolve some disagreements amongst global aerosol models, for instance the difference in volcanic SO_2 forcing efficacy for Pinatubo (see section 3.3.2). The results of this work will help constrain the contribution of stratospheric aerosols to the early 21st century global warming hiatus period, the effects from hypothetical geoengineering schemes, and other climate processes that are influenced by the stratosphere. Overall they provide an excellent opportunity to answer some of these questions as part of the greater WCRP SPARC and CMIP6 efforts. For example, the CMIP6 Geoengineering Model Intercomparison Project (GeoMIP, Kravitz et al., 2015) investigates common ways in which climate models treat various geoengineering scenarios some of them via sulphate aerosols (e.g. Tilmes et al., 2015). However, there is a large inter model spread for the cooling efficiency of sulphate aerosol, i.e. the normalized cooling rate per injected unit of sulphur (Moriyama et al., 2016). ISA-MIP is therefore of special importance for GeoMIP as it could help to understand the reason for these uncertainties, to better constrain the forcing efficiency and to improve future scenarios. Furthermore it is so far not clear whether the large inter-model spread of the CMIP5 models in the simulated post-volcanic climate response mostly depends on uncertainties in the imposed volcanic forcing or on an insufficient representation of climate processes. To discriminate the individual uncertainty factors it is useful to develop standardized experiments/model activities that systematically address specific uncertainty factors. Hence ISA-MIP, which covers the uncertainties in the pathway from the eruption source to the volcanic radiative forcing, will complement the CMIP6 VolMIP project (Zanchettin et al., 2016) which addresses the pathway from the forcing to the climate response and the feedback, by studying the uncertainties in the post-volcanic climate response to a well-defined volcanic forcing. ISA-MIP also complements the chemistry climate model initiative CCMI (Eyring et al., 2013) and the Aerosol Comparison (AeroCom) initiative (Schulz et al., 2006) as well as the Aerosol Chemistry Model Intercomparison Project (AerChemMIP, Collins et al., 2017) as it concentrates on stratospheric aerosol which is not in the focus of all these activities.

As well as identifying areas of agreement and disagreement among the different complexities of models in top-level comparisons focussing on fields such as zonal-mean mid-visible AOD and extinction profiles in different latitudes, ISA-MIP ~~we~~ also intend to explore relationships between key parameters. For example, how does sulphate deposition to the polar ice sheets relate to volcanic forcing in the different interactive stratospheric aerosol models that predict the transport and sedimentation of the particles? Or how do model "spectral extinction curves" evolve through the different volcanically-perturbed periods and how do they relate to simulated effective radius compared to the theoretical approach to derive effective radius from Stothers (1997; 2001). There is considerable potential to apply the model uncertainty analysis to make new statements to inform our confidence of volcanic forcings derived from ice core and astronomical measurements for eruptions before the in-situ measurement era.

Code and data availability

The model output from the all simulations described in this paper will be distributed through the World Data climate Center https://www.dkrz.de/up/systems/wdcch with digital object identifiers (DOIs) as-signed. The model output will be freely accessible through this data portal after registration.

Author contributions.

CT, GWM VA, RH, LAL, AS, CB, SC MC, SSD, TD, JME, MJM, RN, JXS, MT and D.W designed the experiments. CT and GWM coordinated the writing, and drafted the manuscript. All authors have contributed to the writing and have approved of the final version of the manuscript.

Competing interests.

The authors declare that they have no conflict of interest.

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Tables

Experiment	Focus	Number of specific experiments	Years per experiment	Total years [A]	Knowledge-gap to be addressed
Background Stratospheric Aerosol [BG]	Stratospheric sulphur budget in volcanically quiescent conditions	1 mandatory + 2 recommended	20	20(60)	**20 year climatology** to understand **sources and sinks** of stratospheric background aerosol, assessment of sulfate aerosol load under **volcanically quiescent conditions**
Transient Aerosol Record [TAR]	Transient stratospheric aerosol properties over the period 1998 to 2012 using different volcanic emission datasets	4 mandatory +3 optional experiments recommended are 5 (see also Table 4)	15	60 (75,105)	Evaluate models over the period 1998-2012 with different volcanic emission data sets **Understand drivers and mechanisms for observed stratospheric aerosol changes since 1998**
Historic Eruption SO_2 Emission Assessment [HErSEA]	Perturbation to stratospheric aerosol from SO_2 emission appropriate for 1991 Pinatubo, 1982 El Chichón,1963, Agung	for each (x3) eruption (Control, median and 4 (2x2) of hi/lo deep/shallow (see also Table 6)	4 recom. 6	180 (270)	Assess how injected SO_2 propagates through to radiative effects for different historical major tropical eruptions in the different interactive stratospheric aerosol models Use stratospheric aerosol measurements to constrain uncertainties in emissions and gain new observationally-constrained volcanic forcing and surface area density datasets **Explore the relationship between volcanic emission uncertainties and volcanic forcing uncertainties**
Pinatubo Emulation in Multiple Models [PoEMS][B]	Perturbed parameter ensemble of runs to quantify uncertainty in each model's predictions	10 experiments per parameter , where the number of parameters refers to the minimum (3), reduced (5) or standard (8) parameter set (see also Table 10) Each model to vary , 5 or 3 of 8 parameters (7 per parameter = 56 35 or 21)	5 per parameter3 per experiment[C]	90280, 175 or 105(150, 240) (8, 5 or 3)	Intercompare Pinatubo perturbation to strat- aerosol properties with full uncertainty analysis over PPE run by each model. Quantify sensitivity of predicted Pinatubo perturbation stratospheric aerosol properties and radiative effects to uncertainties in injection settings and model processes Quantify and intercompare sources of uncertainty in simulated Pinatubo radiative forcing for the different complexity models.

[A] Each model will need to include an appropriate initialization and spin-up time for each ensemble member (~3-6 years depending on model configuration).

[B] Note, that we are aware that some of the structural parameter variations in PoEMS will introduce some inherent drift in stratospheric aerosol properties for the background control run. However, initial test runs suggest the effect will be much larger for the volcanic perturbation. We therefore expect the effect of the control drift on derived radiative forcings to be small. Models running tropospheric and stratospheric aerosol interactively will need to restrict the parameter scaling to the stratosphere. As explained in the caption to Table 11 and section 3.4, models will need to restrict the PoEMS parameter-scaling to volcanically-enhanced air masses (either via total-sulphur-vmr threshold or passive volcanic SO2 tracer)

C Although the Pinatubo enhancement to the stratospheric aerosol layer remained apparent until 1997 (e.g. Wilson et al., 2008), whereas the HErSEA experiments will continue longer, the PoEMS analysis will require only 3 post-eruption years to be run, as this gives sufficient time after the peak aerosol to characterize decay timescales robustly (e.g. ASAP2006, chapter5) .

Table 1 General overview of the SSIRC ISA-MIP experiments.

Measurement/Platform	Time period 1998-2014	Reference
SO_2 profile/MLS	2004-2011	Pumphrey et al., 2015
SO_2 profile/MIPAS	2002-2012	Höpfner et al., 2013; 2015
Aerosol extinction profile, size/SAGE II	1998-2005	Russell and McCormick, 1989
Aerosol extinction profile, size/OSIRIS	2001-2011	McLinden et al., 2012; Rieger et al., 2015
Aerosol extinction profile/GOMOS	2002-2021	Vanhellemont et al., 2010
Aerosol extinction profile/SCIAMACHY	2002-2012	Taha et al., 2011; von Savigny et al. 2015
Aerosol extinction profile/CALIOP	2006-2011	Vernier et al., 2009, 2011a,b
Aerosol extinction or AOD merged products	1998-2011	Rieger et al., 2015
AOD from AERONET and lidars		Ridley et al., 2014
Surface area density		Kovilakam and Deshler, 2015 Eyring et al. (2013)

Table 2: List of stratospheric aerosol and SO_2 observations available for the BG and TAR time period.

Exp- Name	Specific description / Volcanic emission	Period	Ensemble Size	Years per member	Tier
BG_QBO	Background simulation	Time slice year-2000 monthly-varying with internal or nudged QBO	1	20	1
BG_NQBO	Perpetual easterly phase of the QBO for the whole simulation	Time slice year-2000 monthly varying without QBO	1	20	2
BG_NAT	Only natural sources of aerosol (including biomass burning)	Time slice year-2000 monthly varying with internal of nudged QBO (when possible)	1	20	2

Table 3: Overview of BG experiments.

Volcanic Database	VolcDB1	VolcDB2	VolcDB3	VolcDB4	VolcDBSUB	VolcDB1_3D
Covering period	Dec/1997 - Apr/2012	Jan/1990 - Dec/2014	1978-2014	1979-2010		Dec/1997- Apr/2012
Observational data sets	MIPAS,.GOMOS, SAGEII, TOMS, OMI	OMI, OMPS, IASI, TOMS, GOME/2, , AIRS, MLS, MIPAS	TOMS, HIRS/2, AIRS, OMI, MLS, IASI and OMPS	TOMS, OMI		MIPAS,.GOMOS , SAGEII, TOMS, OMI
Reference	Brühl et al. (2018~~15~~), Bingen et al. (2017), ~~Table S6~~ https://cera-www.dkrz.de/WDCC/ui/cerasearch/entry?acronym=SSIRC_1	Mills et al. (2016, Neely and Schmidt (2016)) http://catalogue.ceda.ac.uk/uuid/bfbd5ec825fa422f9a858b14ae7b2a0d	Carn et al. (2016) https://measures.gesdisc.eosdis.nasa.gov/data/SO2/MSVOLSO2L4.2/)	Diehl et al.,(2012), AeroCom-II HCA0 v1/v2, http://aerocom.met.no/download/emissions/HTAP	Subset of 8 volcanoes Contains SO_2 emissions and plume altitudes averaged over the 3 mandatory databases, details are given in the appendix (Table ~~.~~ S6)	3D netCDF Brühl et al. (2018~~5~~), Bingen et al. (2017), ~~Table S.6~~ https://cera-www.dkrz.de/WDCC/ui/cerasearch/entry?acronym=SSIRC_1

Table 4: Overview of volcanic emission data sets for the different TAR experiments. Sensor acronyms: (MIPAS: Michelson Interferometer for Passive Atmospheric Sounding; GOMOS: Global Ozone Monitoring by Occultation of Stars TOMS: Total Ozone Mapping Spectrometer; OMI: Ozone Monitoring Instrument; OMPS: Ozone Mapping and Profiler Suite; IASI: Infrared Atmospheric Sounding Interferometer; GOME: Global Ozone Monitoring Experiment; AIRS: Atmospheric Infrared Sounder; MLS: Microwave Limb Sounder; HIRS: High-resolution Infrared Radiation Sounder; (References to the observational data and emission sources included are given in the reference paper and for VolcDB1(_3D) also in Table S2.1. VolcDB1_3D is a three-dimensional database, containing the spatial distributions of the injected SO_2 as initially observed by the satellite instruments. In both versions of VolcDB1, the integral SO_2 mass of each injection is consistent.

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Exp- Name	**Volcanic Database Name**	**Specific description**	**Period**	**Years per member**	**TiER**
TAR_base	--	No sporadically erupting volcanic emission	Transient 1998-2012 monthly-varying	15	1
TAR_db1	**VolcDB1**	Volcanic emission data set (Bruehl et al., 2015 and updates)	Transient 1998-2012 monthly-varying	15	1
TAR_db2	**VolcDB2**	Volcanic emission data set (Mills et al. 2016)	Transient 1998-2012 monthly-varying	15	1
TAR_db3	**VolcDB3**	Volcanic emission data set (Carn et al. 2016)	Transient 1998-2012 time-varying	15	1
TAR_db4	**VolcDB4**	Volcanic emission data set (Diehl et al. 2012) and updates	Transient 1998-2010 time-varying	13	3
TAR_sub	**VolcDBSUB**	subset of strongest 8 volcanoes; averaged SO_2 emissions and averaged injection heights from **VolcDB1**/2/3	Transient 1998-2012 monthly-varying	15	2
TAR_db1_3D	**VolcDB1_3D**	**netCDF** version of volcanic emission data set **VolcDB1** (Bruehl et al., 2015 and updates)	Transient 1998-2012 monthly-varying	15	3

Table 5: Overview of TAR experiments.

Exp- Name	Specific description / Volcanic emission	Period	Ensemble Size	Years per member	TiER
HErSEA_Pin_Em_Ism	Pinatubo episode, SO_2 Emission = medium, Inject shallow @medium-alt.	Transient **1991-1995** incl. GHGs & ODSs (monthly-varying SST & sea-ice from HadISST as for CCMI)	3	5	1
HErSEA_Pin_Eh_Ism	Pinatubo episode, SO_2 Emission = high, Inject shallow @medium-alt.		3	5	1
HErSEA_Pin_El_Ism	Pinatubo episode, SO_2 Emission = low, Inject shallow @medium-alt		3	5	1
HErSEA_Pin_Em_Isl	Pinatubo episode, SO_2 Emission = medium, Inject shallow @low-alt		3	5	2
HErSEA_Pin_Em_Idp	Pinatubo episode, SO_2 Emission= medium, Inject over deep altitude-range		3	5	2
HErSEA_Pin_Cntrol	Pinatubo episode, No Pinatubo SO_2 emission		3	5	1
HErSEA_ElC_Em_Ism	El Chichón episode, SO_2 Emission= medium, Inject shallow@ medium-alt	Transient 1982-1986 incl. GHGs & ODSs (monthly-varying SST and sea-ice from HadISST as for CCMI)	3	5	1
HErSEA_ElC_Eh_Ism	El Chichón episode, SO_2 Emission= high, Inject shallow@medium-alt		3	5	1
HErSEA_ElC_El_Ism	El Chichón episode, SO_2 Emission = low, Inject shallow@medium-alt		3	5	1
HErSEA_ElC_Em_Isl	El Chichón episode, SO_2 Emission=medium, Inject shallow@low-altitude		3	5	2
HErSEA_ElC_Em_Idp	El Chichón episode, SO_2 Emission= medium, Inject over deep altitude-range		3	5	2
HErSEA_ElC_Cntrol	El Chichón episode no El Chichón SO_2 emission		3	5	1
HErSEA_Agg_Em_Ism	Agung episode SO_2 Emission= medium, Inject shallow @medium-alt	Transient 1963-1967 incl. GHGs & ODSs(monthly-varying SST and sea-ice from HadISST as for CCMI)	3	5	1
HErSEA_Agg_Eh_Ism	Agung episode, SO_2 Emission= high, Inject shallow @medium-alt		3	5	1
HErSEA_Agg_El_Ism	Agung episode, SO_2 Emission = low, Inject shallow @medium-alt		3	5	1
HErSEA_Agg_Em_Isl	Agung episode, SO_2 Emission = medium, Inject shallow @low-alt		3	5	2
HErSEA_Agg_Em_Idp	Agung episode, SO_2 Emission =medium, Inject over deep altitude-range		3	5	2
HErSEA_Agg_Cntrol	Agung episode no Agung SO_2 emission		3	5	1

Table 6: Overview of HErSEA experiments

Eruption	Measurement/platform	References
Pinatubo	Extinction/AOD [multi-l]: SAGE-II, AVHRR, HALOE,CLAES Balloon-borne size-resolved concentration profiles (CPC, OPC) Impactors on ER2 (AASE2), FCAS and FSSP on ER2 (AASE2) Ground-based lidar; airborne lidar Ship-borne lidar measurements	Hamill and Brogniez (SPARC, 2006, and references therein) Deshler et al (1994, Kiruna, EASOE), Deshler et al. (2003) Pueschel et al. (1994), Wilson et al. (1993), Brock et al. (1993) NDACC archive; Young, S. A et al. (1994), Browell et al., (1993) Avdyushin et al. (1993); Nardi et al. (1993), Stevens et al. (1994)
El-Chichón	Satellite extinction/AOD 1000nm (SAM-II) Balloon-borne particle concentration profiles Ground-based lidar	Hamill and Brogniez (SPARC, 2006 & references therein) Hofmann and Rosen (1983; 1987). NDACC archive
Agung	Surface radiation measurements (global dataset gathered in Dyer and Hicks; 1968) Balloon-borne measurements Ground-based lidar, searchlight and twilight measurements Aircraft measurements	Dyer and Hicks (1965), Pueschel et al. (1972), Moreno and Stock (1964), Flowers and Viebrock (1965) Rosen (1964; 1966, 1968), Pittock (1966) Clemesha et al. (1966), Grams & Fiocco (1967), Kent et al. (1967) Elterman et al., (1969), Volz (1964; 1965; 1970) Mossop et al. (1963; 1964), Friend (1966)

Table 7 List of stratospheric aerosol observation datasets from the 3 large eruptions of the 21st century (Agung, El Chichón and Mt. Pinatubo). For NDACC archive, see http://www.ndsc.ncep.noaa.gov/data/

Eruption	Location	Date	SO_2 (Tg)	Shallow x 2	Deep
Mt. Pinatubo	15°N,120°E	15/06/1991	10-20 (14)	18-20,21-23km	18-25km
El Chichón	17°N,93°W	04/04/1982	5-10 (7)	22-24,24-26km	22-27km
Mt. Agung	8°S,115°E	17/03/1963	5-10 (7.)	17-19,20-22km	17-23km

Table 8: Settings to use for initialising the mini-ensemble of interactive stratospheric aerosol simulations for each eruption in the HErSEA experiment. For Pinatubo the upper range of SO_2 emission is based on TOMS/TOVS SO_2 observations (Guo et al., 2004a). The SO_2 emissions flux ranges and central-values (in parentheses) are specifically for application in interactive stratospheric aerosol (ISA) models, rather than any new data compilation. the lower range and the central values according to some recent Pinatubo studies (Dhomse et al., 2014; Mills et al., 2016; Sheng et al., 2015a) which have identified a modest downward-adjustment of initial observed SO_2 amounts to agree to HIRS/ISAMS measurements of peak sulphate aerosol loading (Baran and Foot, 1994). The adjustment assumes either uncertainties in the satellite measurements or that loss pathways in the first few weeks after these eruptions are either underpredicted (e.g. due to coarse spatial resolution) or omitted completely (accommodation onto ash/ice) in the ISA models. The El Chichón SO_2 central estimate is taken from Krueger et al. (2008), and an emission range based on assumed ±33% while for Agung the SO_2 emission estimate is from Self and King (1996). For Pinatubo, injection height-ranges for the two shallow and one deep realisation are taken from Antuña et al. (2002). The El Chichón values are based on the tropical lidar signal from Figure 4.34 of Hamill and Brogniez (2006), whereas for Agung we considered the measurements presented in Dyer and Hicks (1968) including balloon soundings (Rosen, 1964) and ground-based lidar (Grams and Fiocco, 1967).

SO_2 mass (Tg S)	Study	SO_2 Height (km)
5	Dhomse et al., 2014	19-27
5	Mills et al. (2016)	18-20
7	Sheng et al. (2015a;b)	17-30
8.5	Timmreck et al. (1999a;b)	20-27
8.5	Niemeier et al. (2009); Toohey et al. (2011)	24
8.5	Brühl et al., (2015)	18-26*
10	Pitari and Mancini (2002)	18-25
10	Oman et al. (2006)	19-29
10	Aquila et al. (2012; 2013)	16-18, 17-27
10	English et al. (2013)	15.1-28.5

Table 9: List of SO_2 injection settings used in different interactive stratospheric aerosol model simulations of the 1991 Mount Pinatubo eruption. * main peak at 23.5km, secondary peak at 21km.

	Parameters	Minimum set	Reduced set	Standard set	Uncertainty range
1	Injected SO_2 mass	X	X	X	5 Tg-S – 10 Tg-S
2	Mid-point height of 3km-thick injection	X	X	X	18km – 30km
3	Latitudinal extent of the injection	X	X	X	Factor 0-1 to vary from 1-box injection at 15N (factor=0) to equator-to-15N (factor=1) *
4	Sedimentation velocity		X	X	Multiply model calculated velocity by a factor 0.5 to 2.
5	SO_2 oxidation scaling		X	X	Scale gas phase oxidation of SO_2 by a factor 0.5 to 2
6	Nucleation rate of sulfate particles			X	Scale model calculated rate by a factor 0.5 to 2.
7	Sub-grid particle formation factor.			X	Emit fraction of SO_2 as sulphuric acid particles formed at sub-grid-scale (0 to 10%)
8	Coagulation rate			X	Scale the model calculated rate by a factor 0.5 to 2.

Table 10: Groups will need to translate the 0-1 latitude-spread parameter into a sequence of fractional injections into all grid boxes between the equator and 15 °N. For example for a model with 2.5 degree latitude resolution, the relative injection in the 6 latitude bins between 0 and 15N would take the form [0,0,0,0,0,0,1] for extent factor=0, and [0.167,0.167, 0.167,0.167, 0.167,0.167] for extent factor=1. Injection ratios for intermediate values of the spread factor would be calculated by interpolation between these two end member cases.

Exp- Name	Specific description / Volcanic emission	Period	TIER
PoEMS_OAT_med	SO_2 Emission = medium, Inject shallow @medium-alt. Processes unperturbed.		1
PoEMS_OAT_P4h	SO_2 Emission = medium, Inject shallow @medium-alt. Sedimentation rates doubled		2
PoEMS_OAT_P4l	SO_2 Emission = medium, Inject shallow @medium-alt. Sedimentation rates halved		2
PoEMS_OAT_P5h	SO_2 Emission = medium, Inject shallow @medium-alt. SO_2 oxidation rates doubled		3
PoEMS_OAT_P5l	SO_2 Emission = medium, Inject shallow @medium-alt. SO_2 oxidation rates halved		3
PoEMS_OAT_P6h	SO_2 Emission = medium, Inject shallow @medium-alt. Nucleation rates doubled	Transient 1991-1995	3
PoEMS_OAT_P6l	SO_2 Emission = medium, Inject shallow @medium-alt. Nucleation rates halved		3
PoEMS_OAT_P7h	SO_2 Emission = medium, Inject shallow @medium-alt. % SO_2 as primary SO_4 x2		3
PoEMS_OAT_P7l	SO2 Emission = medium, Inject shallow @medium-alt. % SO_2 as primary SO_4 x0.5		3
PoEMS_OAT_P8h	SO_2 Emission = medium, Inject shallow @medium-alt. Coagulation rates doubled		2
PoEMS_OAT_P8l	SO_2 Emission = medium, Inject shallow @medium-alt. Coagulation rates halved		2

Table 11: Overview of PoEMS One-At-a-Time" (OAT) test runs. Note that when imposing the parameter-scaling, the models should only enact the change in volcanically-enhanced air masses (where the total sulphur volume mixing ratio exceeds a threshold suitable for their model). Perturbing only the volcanically-enhanced air masses will ensure, pre-eruption conditions and tropospheric aerosol properties remains unchanged by the scalings.

Figures

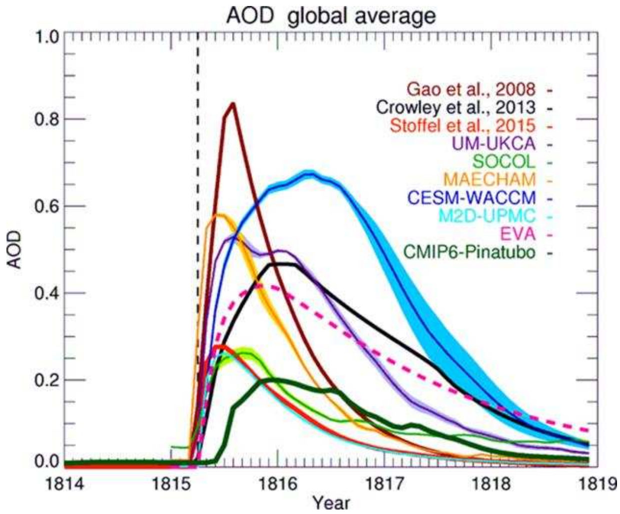

Figure 1 Uncertainty in estimates of radiative forcing parameters for the 1815 eruption of Mt. Tambora: Global-average aerosol optical depth (AOD) in the visible band from an ensemble of simulations with chemistry–climate models forced with a 60 Tg SO_2 equatorial eruption, from the Easy Volcanic Aerosol (EVA, Toohey et al., 2016b) module with 56.2 Tg SO_2 equatorial eruptions (magenta thick dashed line), from Stoffel et al. (2015), from Crowley and Unterman (2013), and from Gao et al. (2008, aligned so that the eruption starts on April 1815). The estimate for the Pinatubo eruption as used in the CMIP6 historical experiment is also reported for comparison. The black triangle shows latitudinal position and timing of the eruption. Chemistry–climate models are CESM (WACCM) (Mills et al., 2016), MAECHAM5-HAM (Niemeier et al., 2009), SOCOL (Sheng et al., 2015a), UM-UKCA (Dhomse et al., 2014), and CAMB-UPMC-M2D (Bekki, 1995; Bekki et al., 1996). For models producing an ensemble of simulations, the line and shading are the ensemble mean and ensemble standard deviation respectively. Figure from Zanchettin et al. (2016).

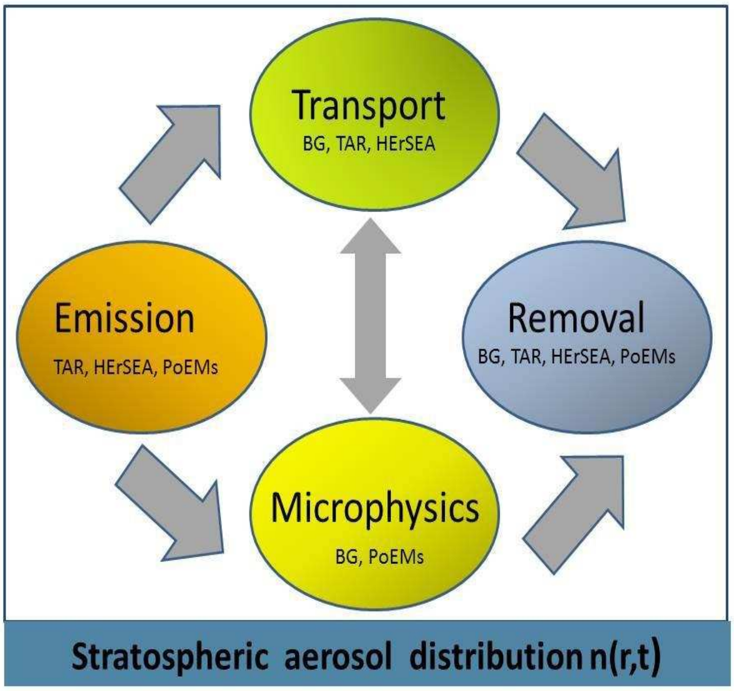

Figure 2 Schematic overview over the processes that influence the stratospheric aerosol size distribution. The related SSiRC experiments are listed below. BG stands for "BackGround", TAR for "Transient Aerosol Record", HErSEA for "Historical Eruption SO_2 Emission Assessment"andPoEMs for "Pinatubo Emulation in Multiple models".

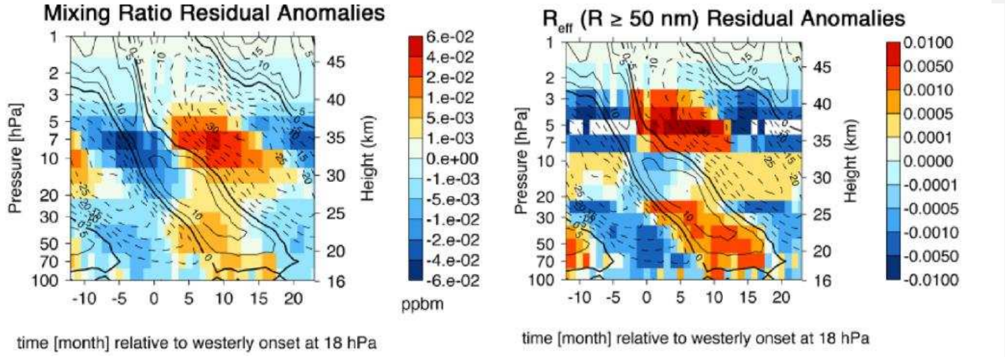

Figure 3. (a) Composite of QBO-induced residual anomalies in the MAECHAM5-SAM2 modelled aerosol mass mixing ratio with respect to the time of onset of westerly zonal mean zonal wind at 18 hPa. Black contours denote the residual zonal wind. Dashed lines represent easterlies, contour interval is 5ms (b) same but for the modelled effective radius of aerosols with R≥50 nm. Figure from Hommel et al. (2015).

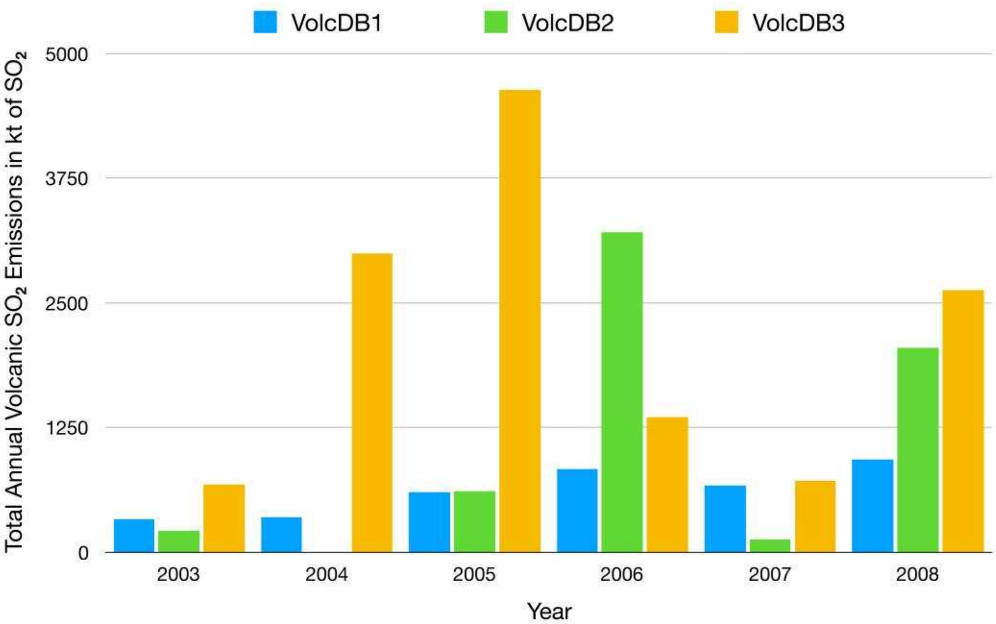

Figure 4: Annual total volcanic sulfur dioxide (SO_2) emission from three different emission data sets between 2003 and 2008 to be used in the TIER1 MITAR experiments. VolcDB1 (Bingen et al., 2017) considers only stratospheric SO_2 emissions, VolcDB2(Neely and Schmidt, 2016) and VolcDB3 (Carn et al., 2016) consider both tropospheric and stratospheric SO_2 emission.

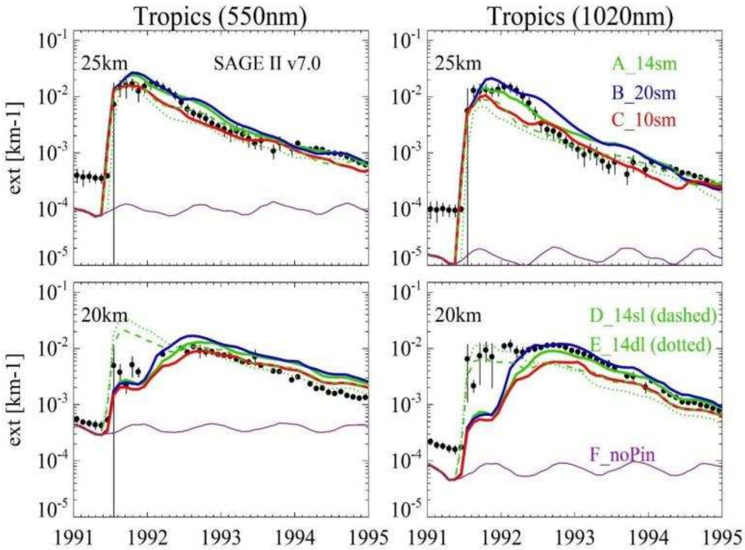

Figure 5: Example results from interactive stratospheric aerosol simulations with the UM-UKCA model (Dhomse et al., 2014) of 5 different SO_2-injection-realisations of the 1991 Pinatubo eruption (see Table 3.3.1), The model tropical –mean extinction in the mid-visible (550nm) and near-infra-red (1020nm) is compared to that from SAGE-II measurements. Only 2 of the 5 injection realisations inject below 20km and the impact on the timing of the peak, and general evolution of the aerosol optical properties is apparent. In this model the growth to larger particle sizes and subsequent sedimentation to lower altitudes is able to explain certain signatures seen in the satellite data (see also Mann et al., 2015).

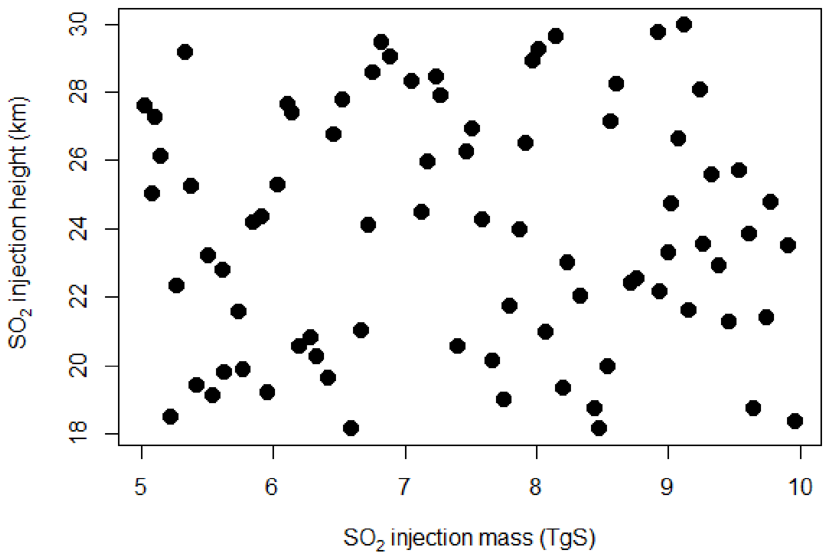

Figure 6 Illustration of the latin hypercube sampling method. Each dot represents the value used in one of the particular simulations with a perturbed parameter ensemble (PPE) with 50 members (realisations/integrations).

List of Abbreviations

AEROCOM	Aerosol Comparisons between Observations and Models
AOD	Aerosol Optical Depth
AMOC	Atlantic Meridional Overturning Circulation
ASAP2006	Assessment of Stratospheric Aerosol properties (WMO, 2006)
AVHRR	Advanced Very High Resolution Radiometer
BDC	Brewer-Dobson Circulation
CALIOP	Cloud-Aerosol Lidar with Orthogonal Polarization
CALIPSO	Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observations
CATS	Cloud-Aerosol Transport System
CCM	Chemistry Climate Model
CCMVal	Chemistry-Climate Model Validation Activity
CCMI	Chemistry-Climate Model Initiative
CCN	Cloud Condensation Nuclei
CDN	Cloud Droplet Number Concentration
CDR	Cloud Droplet Radius
CMIP	Coupled Model Intercomparison Project
CMIP5	Coupled Model Intercomparison Project, phase 5
CMIP6	Coupled Model Intercomparison Project, phase 6
DJF	December-January-February
DWD	Deutscher Wetterdienst
ECHAM	European Center/HAMburg model, atmospheric GCM
EGU	European Geophysical Union
ECMWF	European Centre for Medium-Range Weather Forecasting
EESC	Equivalent Effective Stratospheric Chlorine
ENSO	El Niño Southern Oscillation
ENVISAT	Environmental Satellite
ERA-Interim	ECMWF Interim Re-Analysis
ERBE	Earth Radiation Budget Experiment
ESA	European Space Agency
ESM	Earth System Model
EVA	Easy Volcanic Aerosol
GCM	General Circulation Model
GHG	Green House Gases
GOMOS	Global Ozone Monitoring by Occultation of Stars
HALOE	Halogen Occultation Experiment
HD(CP)2	High definition clouds and precipitation for advancing climate prediction
ISA-MIP	Interactive Stratospheric Aerosol Model Intercomparion Project
ICON	ICOsahedral Nonhydrostatic
IPCC	Intergovernmental Panel on Climate Change
ISCCP	International Satellite Cloud Climatology Project (ISCCP)
ITCZ	Intertropical Convergence Zone
JAXA	Japanese Aerospace Exploration Agency

JJA	June-July-August
LAI	Leaf Area Index
LW	Longwave
LWP	Liquid Water Path
MiKlIP	Mittelfristige Klimaprognosen
MIPAS	Michelson Interferometer for Passive Atmospheric Sounding
MODIS	Moderate Imaging Spectroradiometer
MPI-ESM	Earth System model of Max Planck Institute for Meteorology
NAO	North Atlantic Oscillation
NH	Northern hemisphere
OLR	Outgoing longwave radiation
OMI	Ozone Monitoring Instrument
OMPS	Ozone Mapping and Profiler Suite
OMPS-LP	Ozone Mapping and Profiler Suite–Limb Profiler
OPC	Optical Particle Counter
OSIRIS	Optical Spectrograph and InfraRed Imager System
PDF	Probability Density Function
POAM	Polar Ozone and Aerosol MeasurementPSD
PSD	Particle Size Distribution
QBO	Quasi-biennial oscillation
RF	Radiative Forcing
RH	Relative Humidity
SAOD	Stratospheric Aerosol Optical Depth
SAGE	Stratospheric Aerosol and Gas Experiment
SAM	Southern Annular Mode
SCIAMACHY	Scanning Imaging Absorption Spectrometer for Atmospheric Chartography
SH	Southern Hemisphere
SPARC	Stratosphere-troposphere Processes And their Role in Climate
SSiRC	Stratospheric Sulfur and its Role in Climate
SST	Sea Surface Temperature
SW	Shortwave
TCS	Transient Climate Sensitivity
ToA	Top of the Atmosphere
TOMS	Total Ozone Mapping Spectrometer
TOVS	TIROS Operational Vertical Sounder
VEI	Volcanic Explosivity Index
VolMIP	Model Intercomparison Project on the climate response to Volcanic forcing

Supplementary Information to: The Interactive Stratospheric Aerosol Model Intercomparison Project (ISA-MIP): Motivation and experimental design

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Specifications	Reference
Greenhouse gases ODPs	As recommended for the SPARC CCMI hindcast scenario REF-C1SD (Eyring et al, 2013) http://www.met.reading.ac.uk/ccmi/?page_id=11
SST and SIC	Hadley Centre Sea Ice and Sea Surface Temperature data set (HADISST, Rayner et al., 2003) https://www.metoffice.gov.uk/hadobs/hadisst/

Table S1: Overview of background conditions.

Sulphur emission	Reference
SO_2 Anthropogenic	From MACC-CITY (Granier et al., 2011) for time period considered and as extended back to 1960 on ECCAD website http://eccad.sedoo.fr/eccad_extract_interface/JSF/page_login.jsf
SO_2 Biomass burning	Biomass burning: GFEDv4 (http://www.globalfiredata.org/index.html) From MACC-CITY (Granier et al., 2011) for time period considered and as extended back to 1960 on ECCAD website)
Continuously degassing volcanoes	"continuous_volc.1x1" from Aerocom-I (Dentener et al., 2006) based on Andres and Kasgnoc (1998) which presents an average estimate of the contribution of silent degassing volcanoes to the global sulphur budget http://aerocom.met.no/download/emissions/AEROCOM_B-PRE/other_ascii/
DMS	Sea water concentration from Lana et al. (2011) is recommended https://www.bodc.ac.uk/solas_integration/implementation_products/group1/dms/ Biogenic modeller's choice
OCS	Concentrations are fixed at surface and equal to 510 pptv (Montzka et al., 2013; ASAP2006)

Table S2: Overview of sulphur emission.

Name	Description
nh_50	Passive tracer with fix surface concentration equal to 100 ppb between 30˚N and 50˚N and equal to 0 outside of this latitudinal band, e-folding decay time of 50 days
tr_50	Passive tracer with fix surface concentration equal to 100 ppb between 20˚S and 20˚N and equal to 0 outside of this latitudinal band, e-folding decay time of 50 days;
sh_50	Passive tracer with fix surface concentration equal to 100 ppb between 50˚S and 30˚S and equal to 0 outside of this latitudinal band, e-folding decay time of 50 days.
AOA	Passive tracer for the stratospheric mean age-of-air. Modelling groups can use their existing implementation or implement a tracer with a global fixed surface layer mixing ratio of 0 ppbv and a uniform unspecified fixed source (at all levels) everywhere else, which must be constant in space and time.
ST80_25	Passive tracer to estimate the exchange from the stratosphere to the troposphere. This is achieved by fixing the mixing ratio above 80hPa (200ppbv) to a constant value, and imposing a uniform fixed 25-day exponential decay in the troposphere only.
Volc	Passive volcanic tracer for the HerSEA experiments. The tracer is initialized in the same way as the volcanic SO_2 emission, with an initial value of 1.

Table S3: Suggested passive tracers mostly following the CCM protocol (Eyring et al., 2013).

Long name	Variable name	Unit	Category	Comment
grid-cell area	area	m^2	1	
land fraction	landf	1	1	Please express "X_area_fraction" as the fraction of horizontal area occupied by X.
surface altitude	orog	m	1	"Surface" means the lower boundary of the atmosphere. Altitude is the (geometric) height above the geoid, which is the reference geopotential surface.
Meteorology				
Precipitation	precip	kg m^{-2} s^{-1}	1	Includes all types: rain, snow, large-scale, convective, etc.
surface temperature	tas	K	1	
surface air pressure	ps	Pa	1	"Surface" means the lower boundary of the atmosphere.
Cloud fraction	clt	%	1	Cloud fraction as seen from top or surface
tropopause_air_pressure	ptp	Pa	2	2D monthly mean thermal tropopause calculated using WMO tropopause definition on 3d temperature
tropopause_air_temperature	tatp	K	2	See above
tropopause_altitude	ztp	M	2	See above
Budget				
Load of H2SO4 (aerosol)	loadso4	kg m^{-2}	1	Units of the particle-phase-sulphur should be using mass of H2SO4
Load of SO2(g)	loadso2	kg m^{-2}	1	
Load of H2SO4(g)	loadh2so4	kg m^{-2}	1	
Load of OCS	loadocs	kg m^{-2}	1	
Load of DMS	loaddms	kg m^{-2}	2	
Load of H2S	loadh2s	kg m^{-2}	3	
Load of CS2	loadcs2	kg m^{-2}	3	
Removal				
dry deposition of DMS	drysdms	kg m^{-2} s^{-1}	2	
dry deposition of SO2	dryso2	kg m^{-2} s^{-1}	1	
dry deposition of H2SO4(g)	dryh2so4	kg m^{-2} s^{-1}	1	
dry deposition of H2SO4(p)	dryso4	kg m^{-2} s^{-1}	1	
sedimentation of SO4	sedso4	kg m^{-2} s^{-1}	1	
dry deposition of H2S	dryh2s	kg m^{-2} s^{-1}	2	
dry deposition of C2S	dryc2s	kg m^{-2} s^{-1}	2	
wet deposition of SO2	wetso2	kg m^{-2} s^{-1}	1	
wet deposition of H2SO4(p)	wetso4	kg m^{-2} s^{-1}	1	
wet deposition of DMS	wetdms	kg m-^{2} s^{-1}	2	
wet deposition of C2S	wetc2s	kg m^{-2} s^{-1}	2	
wet deposition of H2S	weth2s	kg m^{-2} s^{-1}	2	
Emission				
total emission of SO2	emiso2	kg m^{-2} s^{-1}	1	
total emission of DMS	emidms	kg m^{-2} s^{-1}	2	
total emission of COS	emicos	kg m-2 s-1	1	If available
total emission of DMS	emih2s	kg m-2 s-1	1	
total emission of CS2	emic2s	kg m-2 s-1	3	
Fluxes				
So2 Flux to the tropopause	flxso2	kg m^{-2} s^{-1}	1	
H2SO4(p)Flux through the tropoause (total)	flxso4t	kg m^{-2} s^{-1}	1	
H2SO4 Flux (tropopause) per size class/modes	flxso4_	kg m^{-2} s^{-1}	3	
Flux H2SO4 (p) > 5nm	flxso4p150	kg m^{-2} s^{-1}	2	
Flux H2SO4 (p) >150nm	flxso4p150	kg m^{-2} s^{-1}	2	
Flux H2SO4 (p) >250nm	flxso4p250	kg m^{-2} s^{-1}	2	
Flux H2SO4 (p) >550nm	flxso4p550	kg m^{-2} s^{-1}	2	
Flux H2SO4 (p) >750nm	flxso4p750	kg m^{-2} s^{-1}	2	
Flux H2SO4 (p) >1000nm	flxso4p1000	kg m^{-2} s^{-1}	2	

Radiation				
AOD@386nm	od386aer	1	2	
AOD@453nm	od453aer	1	2	
AOD@525nm	od525aer	1	1	
AOD@750nm	od750aer	1	2	
AOD@870nm	pd870aer	1	2	
AOD@1020nm	od1020aer	1	1	
AOD@3460nm	od3460aer	1	2	
AOD@5260nm	od5260aer	1	2	
AOD@12660nm	od5260aer	1	2	
Surface downwelling SW radiation	rsds	W m^{-2}	1	
Surface upwelling SW radiation	rsus	W m^{-2}	1	
Surface downwelling LW radiation	rlds	W m^{-2}	1	
Surface upwelling LW radiation	rldus	W m^{-2}	1	
Surface downwelling SW flux clear sky	rsdscs	W m^{-2}	2	
Surface upwelling SW flux clear sky	rsuscs	W m^{-2}	2	
Surface upwelling LW flux clear sky	rldcs	W m^{-2}	2	
Surface diffuse SW flux	rsdsdiff	W m^{-2}	2	
Surface diffuse SW flux clear sky	rsdscsdiff	W m^{-2}	2	
TOA Incident	rst	W m^{-2}	2	
TOA downwelling SW radiation	rsdt	W m^{-2}	1	
TOA downwelling LW radiation	rldt	W m^{-2}	1	
TOA outgoing SW radiation	rsut	W m^{-2}	1	
TOA outgoing SW radiation clear sky	rsutcs	W m^{-2}	2	
TOA outgoing LW radiation	rlut	W m^{-2}	1	
TOA outgoing LW radiation clear sky	rlutcs	W m^{-2}	2	
Total photsynthtically FLUX (PAR)	tphotpar	W m^{-2}	3	
photsynthtically FLUX (PAR)	photpar	W m^{-2}	3	

Table S4: Overview of two-dimensional variables requested for ISA-MIP following mainly the AEROCOM protocols: http://aerocom.met.no/protocol.html. (1) indicates mandatory variables, which are in addition shaded, (2) important variables but not required, (3) values which are nice to have for special diagnostic. Monthly mean output is satisfactory except for the meteorological values, which should be provided in daily resolution.

Long name	Variable name	Unit	Category	Comment
Meteorology				
air temperature	ta	K	1	Air temperature is the bulk temperature of the air, not the surface (skin) temperature.
specific humidity	hus	1	1	Specific means per unit mass. Specific humidity is the mass fraction of water vapor in (moist) air.
air mass	airmass	kg m^{-2}	1	Vertically integrated mass content of air in layer
pressure	pfull	Pa	1	Air pressure on model levels
zonal wind	ua	m/s	1	
meridional wind	va	m/s	1	
vertical wind	wa	m/s	1	
geopotential height	Zg	m	1	
cloud fraction	clt3D	%	2	
cloud optical depth	cod3D	1	2	
aerosol water	mmraerh2o	1	3	
convective updraft mass flux	mcu	kg m-^{2} s^{-1}	3	The atmosphere convective mass flux is the vertical transport of mass for a field of cumulus clouds or thermals, given by the product of air density and vertical velocity. For an area-average, cell_methods should specify whether the average is over all the area or the area of updrafts only.
Sulfur Chemistry				
OCS	vmrocs	1	1	
SO2	vmrso2	1	2	
DMS	vmrdms	1	2	
H2S	vmr h2s	1	3	
H2SO4 (g)	vmrh2so4	1	2	
CS2	vmrcs	1	3	
SO3	vmrso3	1	2	
H2SO4 (p) total)	mmso4r	1	1	Mass mixing ratio of sulphate mass (total)
Mass mixing ratio of sulfate mass in each size class				
H2SO4 (p) > 5nm	mmso4r5	1	2	OPC
H2SO4 (p) >150nm	mmso4r15	1	2	OPC
H2SO4 (p) >250nm	mmso4r25	1	2	OPC
H2SO4 (p) >550nm	mmso4r55	1	2	OPC
H2SO4 (p) >750nm	mmso4r75	1	2	OPC
H2SO4 (p) >1000nm	mmso4r100	1	2	OPC
Microphysical processes				
number formation through nucleation	nucpn	m^{-3} s^{-1}	2	
sedimentation of SO4	sedso4	kg m^{-2} s^{-1}	2	Net downward (out-below minus in-above)
H2SO4 condensation flux	conh2so4	kg m^{-2} s^{-1}	2	Net transfer into the particulate phase
Chemistry				
N2O	vmrn2o	1	3	
OH	vmroh	1	1	
O3	vmro3	1	1	
HNO3	vmrhno3	1	3	
NO	vmrno	1	3	
NO2	vmrno2	1	3	
N2O5	vmrn2o5	1	3	
Bulk parameters				
surface area density	sad	m^2/m^3	1	
effective radius	reff	M	1	
Particle numbers				
N total	conccn	m^{-3}	1	number_concentration_of_ambient_aerosol_in_air
N> 5nm	conc5	m^{-3}	2	CPC
N>150nm	conc150	m^{-3}	2	OPC
N>250nm	conc250	m^{-3}	2	OPC
N>550nm	conc550	m^{-3}	2	OPC
N>750nm	conc750	m^{-3}	2	OPC

N>1000nm	conc1000	m^{-3}	2	OPC
Extinction				
Aerosol extinction @386nm	ec386aer	m^{-1}	2	SAGEII/III, (POAM, shipborne lidar)
Aerosol extinction @440nm	ec440aer	m-^{1}	3	
Aerosol extinction @525nm	ec525aer	m-^{1}	1	SAGE-II
Aerosol extinction @750nm	ec750aer	m^{-1}	2	OSIRIS
Aerosol extinction @870nm	ec870aer	m^{-1}	3	
Aerosol extinction @1020nm	ec1020aer	m^{-1}	1	SAGEII
Aerosol extinction @3460nm	ec3460aer	$m^{-}1$	2	HALOE
Aerosol extinction @5260nm	ec5260aer	m-1	2	HALOE
aerosol extinction @12660nm	ec12660aer	m^{-1}	3	ISAMS
Absorption				
aerosol absorption @386nm	abs386aer	m^{-1}	3	SAGEII/III, (POAM, shipborne lidar)
aerosol absorption@440nm	abs440aer	m^{-1}	3	
aerosol absorption @525nm	abs525aer	m^{-1}	2	SAGE-II
aerosol absorption@750nm	abs750aer	m^{-1}	3	OSIRIS
aerosol absorption @870nm	abs870aer	m^{-1}	3	
aerosol absorption @1020nm	abs1020aer	m^{-1}	2	SAGE-II
aerosol absorption @3460nm	abs3460aer	m^{-1}	3	HALOE
aerosol absorption @5260nm	abs5260aer	m^{-1}	3	HALOE
aerosol absorption @12660nm	abs12660aer	m^{-1}	3	ISAMS
asymmetry factor@525nm	asy525aer	1	1	

Table S5: Overview of three-dimensional variables requested for ISA-MIP following mainly the AEROCOM protocols: http://aerocom.met.no/protocol.html. All 3D data to be provided on either host model vertical levels or preferably (if resources allow) on the reference pressure levels 1000, 925, 850, 700, 600, 500, 400, 300, 250,200, 150, 100, 70, 50, 30, 20 & 10 hPa. If possible also on the additional pressure levels: 7, 5, 3, 2, 1 and 0.4 hPa. (1) indicates mandatory variables, which are in addition shaded, (2) important variables but not required, (3) values which are nice to have for special diagnostic. Monthly mean output is satisfactory except for the meteorological values, which should be provided in daily resolution.

Volcano	Lon	Lat	Time	Min Plume Height (km)	Max Plume Height (km)	Mean SO2 (kt)
Manam	145.04	-4.08	27 Jan 2005	18	24	154.67
Soufriere Hills	297.82	16.72;	19 May 2006;;	19	20	185.33
Rabaul/Tavurvur	152.20	-4.27;	7 Oct 2006	17	18	234.0
Okmok	168.10	53.43	12 Jul 2008	10	16	109.0
Kasatochi	175.50	52.18	7 Aug 2008	10	18	1363.33
Sarychev	153.20	48.09	15 Jun 2009	11	17	965.33
Merapi	110.44	-7.54	4 Nov 2010	14	17	282.67
Nabro	41.70	13.37;	13 Jun 2011	9.7	18	1307.0

Table S6: Overview of VolcDSUB, a subset of volcanic emissions, that were derived based on the average mass of SO_2 emitted using VolcDB1, VolcDB2, and VolcDB3. (http://isamip.eu/fileadmin/user_upload/isamip/volc_sub_v185.dat).

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Volcano or region	Time	Latitude	Longitude	Height	SO2 3D str (kt)
Soufriere Hills	26 Dec 1997	16	-62	16	27
Soufriere Hills	4 Jul 1998	16	-62	16	17*
Manam, Cerro Azul, Nyamuragira	7 Oct 1998	-5, 0, -1	144, -90,30	17,17,16	9, 21, 20
Cameroon	31 Mar 1999	4	10	17	52*
Soufriere Hills+	24 Jul 1999	16	-62	17	19*
Tungurahua+Guagua Pinch	16 Nov 1999	-1,0	-78	17	38
Nyamuragira, Tungurahua	4 Feb 2000	-1,0	30,-78	16	23, 2*
Mayon, Hekla, Vanuatu,Tungurahua	29 Feb 2000	13,64,-16, -1	124,-20, 168,-78	16	39,1,8,35
Ulawun (+ Miyakejima)	26 Sep 2000	-5	150	16-18	59*
Nyamuragira	13 Feb 2001	-1	30	16	50*
Ulawun	29 Apr 2001	-5	150	16	51
Mayon, Lopevi	23 Jun 2001	13, -16	124, 168	16	62, 29
Tungurahua, Soufriere Hill	7 Aug 2001	0, 16	-78, -62	16	9, 6
Manam, Nyiragongo	14 Jan 2002	-5, -1	144, 30	17, 15	21, 12

Table S6: Supplement to VolcDB1, extension of table 7 (Bingen et al., 2017) Volcanic SO_2 injections into the stratosphere, derived from MIPAS and OMI/TOMS (Brühl et al., 2015). Updated on the basis of GOMOS, SAGE II(V7.00) and new MIPAS data (from UTLS mode). SO_2 masses above 14km in low latitudes, above 13km in midlatitudes and above 12km in high latitudes. Listed altitudes and latitudes refer to the maxima in MIPAS and SAGE 'plumes'. The given time refers to the center of the first MIPAS 5 day period selected and not the beginning of the eruption. * above 15km.

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