Interactive comment on "ORCHIDEE-PEAT (revision 4596), a model for northern peatland CO2, water and energy fluxes on daily to annual scales" by Chunjing Qiu et al.

We thank the two anonymous referees very much for their constructive comments. In the following, please find our response to the comments. Our responses are in bold, modifications done in the revised manuscript are in blue. All figure and table numbers, line numbers and pages refer to the initial manuscript version.

Referee #1

This manuscript describes a new peatland model implemented in the ORCHIDEE land model. The model was evaluated by comparing modeled water table, LE, GPP, and NEE to measured eddy covariance fluxes from several peatland field sites. The paper is generally well written and the key processes of the model are clearly described. The introduction section includes a useful review of recent peatland models that does a good job of setting the stage for this model. The paper generally does a good job of identifying uncertainties and potential weaknesses in the model that could be addressed in future work, although I think there is some room for improvement in describing some of these issues in more depth.

I think there are a couple of general areas in which the manuscript could be improved:

1. The key peatland-specific changes to the model are focused on peat carbon pools and hydrology, including a new architecture for simulating peat decomposition using acrotelm and catotelm layers. The modifications to plant processes are less dramatic. In my understanding the model uses an existing C3 grass plant functional type and does not introduce any new peatland-specific vegetation processes. Given the focus of model process changes on decomposition rather than plant processes, it seems strange that the evaluation is so focused on GPP. Why not show and evaluate modeled ecosystem respiration instead of or in addition to GPP? Analyzing respiration fluxes would allow a much better evaluation of the key new model features that are specific to peatland processes. Without an evaluation specific to these new processes, it feels like there is a big piece missing.

While our initial focus was on peatland productivity and carbon intake, of course GPP and ER are linked. So we followed the reviewer's suggestion to incorporate an analysis of ecosystem respiration. To do so, we added analyses and discussion of simulated vs. measured ecosystem respiration. In the first set of simulation (S1) in which the modeled water table were used in the carbon module, with the site-specific V_{cmax} , the model showed good performance in capturing both spatial and temporal variations in ER, with r^2 of 0.78, 0.89, 0.86 for daily variations, across-sites annual variations and seasonal variations, respectively, and MEF of 0.75, 0.79, and 0.86, respectively. These results were compared with simulations using a fixed V_{cmax} (the mean of the optimized V_{cmax} , 40 µmol m⁻² s⁻¹), as suggested by the review in the second comment. We conclude that by

taking site-to-site variations in V_{cmax} into consideration, model performances for carbon fluxes (GPP, ER and NEE) were improved. Table4 and Figure4 have been updated to include results of ER and results with the mean of the optimized V_{cmax}, and the description of the results from Line474 to Line487 was rephrased as: "For the 22 sites where NEE and ER measurements were available, the errors in the three carbon fluxes – GPP, ER, NEE were significantly reduced by optimizing V_{cmax} at each site (Table 4, Fig. 4, Fig. S4). With site-specific V_{cmax} values (Site-by-site model performances are shown in Fig. S5 to S10 in Supplementary Materials), the overall (all the daily data from all the 22 sites) performance of the model was good for GPP ($r^2 = 0.76$, MEF = 0.76), ER ($r^2 = 0.78$, MEF = 0.75), and acceptable for NEE ($r^2 = 0.38$, MEF = 0.26) (Fig. 4, Table 4). Seasonal variations in carbon fluxes were well captured by the model ($r^2 = 0.61$ to 0.86). The spatial across-sites gradients of annual mean GPP and ER were generally good, with r^2 of 0.93 and 0.89, and lower for NEE ($r^2 = 0.27$). Compared to simulations with a fixed V_{cmax} (the mean of the optimized values of 40 µmol m⁻² s⁻¹), there were large improvements in capturing spatial gradients of carbon fluxes (r² increased from 0.20 to 0.93, from 0.27 to 0.89 and from 0.16 to 0.27 for GPP, ER and NEE, respectively, while the RMSD reduced by 63%, 48%, and 9%). This result indicates that model-data disagreement can be largely reduced by using site-specific V_{cmax} instead of a fixed (mean) value. In future regional simulations, spatial variations in V_{cmax} should be taken into account. There was, however, no significant improvement in LE, H and WT by using site-specific V_{cmax} values (Table4).". We also compared simulated ER of S1 with the second set of simulation (S2, in which the measured water table was used) with the ER observations: the model showed only a small improvement in reproducing ER when WT_{obs} was used (Table 5 was added to show the results). Fig.S6 was added to show simulated vs. measured ER at each site.

Table 5. Model performance measures of ER simulations for the site-by-site comparison, the comparison across sites, mean seasonal cycle and anomalies, using modeled (S1) and observed (S2) water table (WT). SDSD and LCS are two signals discriminated from the mean squared deviation, see Sect. 3.4.

	Modeled WT used (S1)					Observed WT used (S2)				
Site	RMSD	SDSD	LCS	r^2	MEF	RMSD	SDSD	LCS	r^2	MEF
CZ-Wet	1.45	0.86	0.87	0.81	0.68	1.51	1.05	0.79	0.81	0.66
DE-Bou	0.78	0.03	0.50	0.69	0.64	0.77	0.03	0.50	0.69	0.65
DE-Sfn	0.96	0.10	0.79	0.61	0.59	0.97	0.09	0.82	0.60	0.58
FI-Lom	0.46	0.00	0.19	0.85	0.84	0.45	0.02	0.18	0.85	0.84
IE-Kil	0.44	0.01	0.01	0.09	0.51	0.42	0.01	0.01	0.13	0.48
SE-Deg	0.69	0.26	0.19	0.75	0.62	0.64	0.16	0.23	0.75	0.68
SE-Faj	0.58	0.07	0.08	0.87	0.60	0.59	0.08	0.07	0.88	0.59
US-Los	0.63	0.01	0.39	0.85	0.85	0.60	0.00	0.35	0.87	0.87
Overall	0.79	0.09	0.51	0.78	0.76	0.79	0.09	0.51	0.78	0.76
Across	0.21	0.01	0.06	0.02	076	0.22	0.01	0.06	0.02	0.74
sites	0.51	0.01	0.06	0.82	0.70	0.32	0.01	0.00	0.82	0.74
Seasonal	0.45	0.06	0.15	0.91	0.89	0.44	0.07	0.13	0.92	0.89
Anomalies	0.62	0.07	0.31	0.21	0.19	0.63	0.08	0.31	0.20	0.17

2. The approach to optimizing Vcmax is problematic. The optimized site-specific values are compared to a default value that is well outside the range of values that seem to be appropriate for these sites (within the model at least). Figure S3 demonstrates this very clearly for GPP and NEE: the model using the default Vcmax is not even close to reproducing the observed magnitude of photosynthesis at these sites. As a result, the comparison between optimized and default Vcmax simulations is not very informative. It would be more useful if that comparison used the mean or median of the optimized Vcmax values (which is actually used for a different analysis later in the paper). In that case, it would be possible to evaluate whether site-to-site variations in Vcmax were necessary for improving model fidelity. It's not very informative to show that optimized Vcmax is better than a Vcmax that is much too low for every site.

The fact that the default Vcmax based on observations does not work within the model raises further questions. The paper addresses this very briefly (lines 249-251) but I think a more detailed discussion of why the model Vcmax needs to be so much higher than observations would be useful. Were the other photosynthesis-related parameters (LAI, light absorption, etc) in the model consistent with site measurements? Site-specific optimization of Vcmax could mask other issues with the model, for example underestimates of plant biomass or LAI. I think it would be really helpful to show how modeled LAI compares to measurements, especially among different sites, and whether errors in modeled LAI can explain the latitude/temperature relationship in optimized site Vcmax.

The reviewer raises a fair point that a comparison between the optimized and the default V_{cmax} value (16 µmol m⁻² s⁻¹) is not as informative as it could be in this study. The default value applied by Largeron et al. (2017, Geosci. Model Dev. Discuss.) was derived for three low productivity sites. When this value was applied at our dataset, GPP and NEE were underestimated. Thus, to make a more apples-to-apples test, we added a comparison between optimized and the mean of the optimized V_{cmax} values (40 µmol m⁻² s⁻¹), as suggested by the reviewer. The comparison to the default V_{cmax} is removed from the manuscript. The description of the results from Line474 to Line487 was rephrased, as it is mentioned in our response to the first comment of the reviewer.

Our use of site optimized V_{cmax} is one way to account for large variance in a key ecosystem parameter. There is a large reported variation of V_{cmax} in observations. For instance, V_{cmax} value for Sphagnum at the Old Black Spruce site in Canada were 5, 14 and 6 µmol m⁻² s⁻¹ during spring, summer and autumn respectively, while that for Pleurozium were 7, 5, and 7 µmol m⁻² s⁻¹ (Williams and Flanagan, 1998, PCE); Bubier et al. (2011, Oecologia) reported that V_{cmax} for three ericaceous shrubs (*Vaccinium myrtilloides, Ledum groenlandicum* and *Chamaedaphne calyculata*) at Mer Bleue bog in Canada ranged from 67 to 137 µmol m⁻² s⁻¹ among the control and four nutrient addition treatments (measured V_{cmax} for the three shrubs in the control plots are 84.6 ± 13.5 µmol m⁻² s⁻¹, 78.1 ± 13.4 µmol m⁻² s⁻¹, and 132.1 ± 31.2 µmol m⁻² s⁻¹, respectively); The V_{cmax} value applied by the McGill wetland model for evergreen shrubs is 17 µmol m⁻² s⁻¹, which is the median value of over 50 measurements for *Chamaedaphne calyculata* and *Ledum groenlandicum* (St-Hilaire et al., 2010, Biogeosciences). Wu et al. (2016, Geosci. Model Dev.) used values of 60, 50, 40µmol m⁻² s⁻¹ for evergreen shrubs,

deciduous shrubs and sedges respectively. The optimized model V_{cmax} values in our study ranged from 19 to 89 µmol m⁻² s⁻¹ (the mean value is 40 µmol m⁻² s⁻¹), considering that the model optimized value represents an average for the ecosystem, we argue that the model value is not substantially above observations or values used in other land surface models.

We agree with the reviewer that site-specific optimization of V_{cmax} could compensate for biases in LAI, plant biomass, etc. Unfortunately, at most of the sites, LAI was measured or estimated (by optical in-situ methods, annual litter collection, or from remote sensing) only once during the periods in question. We have an available time-series of measured LAI at IE-Kil – see Fig.S1(a), LAI was overestimated by the model at IE-Kil. Fig.S1 (b) showed that LAI was overestimated at sites with low reported (measured or estimated) LAI and underestimated at sites with higher reported values. As for aboveground biomass, there was no systematic error among sites. We emphasized the bias in LAI in the text, on Page16, Line463: "....., with a mean value of 40 µmol m⁻² s⁻¹. The calibration of V_{cmax} may compensate for biases in other model parameters. A brief comparison between simulated and reported (measured/estimated) LAI and aboveground biomass showed that there are no systematic errors (Fig. S1).".



Fig. S1. (a) Simulated vs. measured leaf area index (LAI) at the blanket bog IE-Kil, Ireland. (b) Simulated vs. reported (measured/estimated) LAI across peatland sites, dashed line is a hypothetical 1:1 regression line. Note that in (b), the reported LAI was estimated at some sites. (c) Simulated vs. measured aboveground biomass, assuming that the carbon content of dry biomass is 50%.

Specific comments:

Lines 173-176: I'm not sure it's that novel that this model is built into a land surface scheme that conserved water, carbon, and energy. Doesn't the LPJ-GUESS model described above have a similar purpose? In any case, if there is not already a peatland submodel built into ORCHIDEE then I wouldn't be that concerned about justifying the purpose of this effort. I think it's clearly valuable to build and evaluate a working peatland submodel within ORCHIDEE.

The reviewer is right, the LPJ-GUESS does describe a similar development, however, there is no water input from surrounding areas (Chaudhary et al., 2016, Biogeosciences), so conservation is scale-dependent. We rephrased the sentences on Page6, Line173 as follows : "This new peat model is incorporated consistently into the land surface scheme in order to conserve water, carbon and energy at scales going from local sites to grid-based

large-scale applications in an Earth System Modeling context."

Line 232: Not all peatlands are grassy. Does this assumption cause issues when applying the model to shrubby or forested peatlands (such as the Old Black Spruce site mentioned a few lines after this)? Were all the peatland sites used for evaluation grassy peatlands?

The sites used for evaluation include grassy, shrubby, and forested peatlands (Table 2). We note the possible discrepancies between model output and observations in the text as suggested by Reviewer#2. Please refer to our response to the third comment of Reviewer#2 (Lines 231-232).

Line 249-251: It's great that the paper brings up this issue of compensating errors, but it would be better if there were some evaluation of whether the model has systematic errors in LAI, etc.

As shown in Fig. S1b, LAI was overestimated at sites with low reported LAI and underestimated at sites with high reported values, there was no systematic error in LAI.

Line 257: "drainage flux reduced to zero": So there is no water flow out of the peatland unless it is flooded? This seems inconsistent with a lot of real peatland systems.

We would like to note that although we considered deep drainage from peatland as negligible due to the low permeability of the catotelm (Ingram et al., 1978, EJSS; Rezanezhad et al., 2016, Chem. Geol.), the waterflow out of the peatland (as runoff) occurs not only when the peatland is flooded. In ORCHIDEE, the partitioning between water infiltration and surface runoff is computed through a time-splitting procedure (d'Orgeval, 2006, PhD thesis), with the maximum infiltration rates described as an exponential probability density distribution. The infiltration-excess water creates runoff. Thus in the model, the infiltration excess water will first fills the above-surface water reservoir, and then leaves the grid cell as runoff.

To clarify this, we added these sentences in the text, Page9, Line259: ".....an above surface water reservoir with a maximum height of 10 cm was added (Fig. 1b). In the model, the partitioning between water infiltration and surface runoff is computed through a time-splitting procedure, with the maximum infiltration rates described as an exponential probability density distribution (d'Orgeval, 2006, Diss. Paris). The infiltration-excess water of peatland first fills the above-surface water reservoir, then leaves the grid cell as runoff. Water in this above-surface reservoir re-infiltrates into the peat soil on the next time step (Largeron et al., 2017).".

Line 299-301 and Fig. S1: The difference between the soil carbon dynamics and the peat carbon dynamics is confusing. Do the peat pools contain the Active/Slow/Passive soil carbon pools, or do they replace them? Fig. S1 suggests that all of these pools are present in the peatland (metabolic litter, structural litter, acrotelm, catotelm, active, slow, passive) but this doesn't seem consistent with the description in the text. If the peat layers are actually replacing the active/slow/passive pools, then Fig. S1 and the text should make that clearer. **The reviewer is right, the description of the carbon module is not clear enough. We improved the description in the text on Page10, Line295 to "……Decomposed litter carbon**

from these two pools is then distributed into three soil carbon pools: the active, slow and passive pool, similar to the CENTURY model (Parton et al., 1988). Both temperature and moisture functions are used to control soil carbon decomposition rates (Text S1). In ORCHIDEE-PEAT, these standard processes are kept the same as in Krinner et al. (2005) for non-peatland vegetation (Fig. S2, black dashed box). For the peatland vegetation, we added a peat carbon module, in which the three soil carbon pools (active, slow, passive) are replaced by two pools forming distinct layers, following Kleinen et al. (2012) (Fig. S2, red dashed box)." and we modified Fig. S2 as follows:



Fig. S2. Schematic overview of litter and soil carbon dynamics in ORCHIDEE-PEAT. For non-peatland vegetation (the black dashed box), decompositions of carbon in the two litter pools and three soil pools, and carbon flows between them are adapted from the CENTURY model (Parton et al., 1988); for peatland vegetation (the red dashed box), the active, slow and passive soil carbon pools are replaced by a two-layered model, following Kleinen et al. (2012).

Line 308-310: Did this use the observed or simulated water table? How would this be handled in larger-scale or global simulations?

The simulated mean summer minimum water table position (WT_{min}) over the observational period is used here. WT_{min} was derived from a 'preparation run (S0)'. Specifically, we first ran the model at each site using the same simulation protocol as described in Sect. 3.3, but with peat carbon module deactivated. Then WT_{min} can be diagnosed from the output of this simulation (S0) and be fed into the model in S1 and S2.

We explained this procedure in the text on Page10, Line308 to: ".....we used the average of simulated minimum summer water table position (WT_{min}) over the observational period to demarcate the boundary between the acrotelm and the catotelm at each site to take into

account local site conditions. We conducted a "preparation run (S0)", in which the model was run at each site using the same protocol (Sect. 3.3), but with the peat carbon module deactivated. WT_{min} was diagnosed from the output of S0 before feeding into the peat carbon module in S1 and S2 (Sect. 3.3). Soil carbon exerts no feedback effects on the soil temperature and hydraulics in the structure of our model, thus S0 and S1 produce the same simulated water table.". In large-scale or global simulations, we can either conduct the same "preparation run" or set WT_{min} to a constant value, for example, Wania et al. (2009, Global Biogeochem. Cycles) and Spahni et al. (2013, Clim. Past.) used 0.3 m as the interface between the acrotelm and the catotelm.

Line 315-316: It would help to show the equation for beta instead of just describing it. Equations for acrotelm height and catotelm depth should also be included. Is the depth of catotelm and total peat depth calculated? What does the model do if water table goes below the bottom of the peat layer? Can it represent a situation with no catotelm layer? Is there mineral soil beneath the bottom of the peat layers?

We added equations of beta (Eq. 9) and acrotelm depth (Eq.10) on Page11, Line325:

$$\beta = \begin{cases} \beta = 1, & WT_{min} - WT \le 0\\ \beta = \frac{H_A - (WT_{min} - WT)}{H_A}, & 0 < WT_{min} - WT < H_A\\ \beta = 0, & WT_{min} - WT \ge H_A \end{cases},$$
(9)

$$H_{A} = \frac{C_{A}}{\rho_{A} \cdot C_{f,A}} , \qquad (10)$$

The depth of catotelm can be calculated using carbon fraction in the catotelm and the catotelm density, as in Kleinen et al. (2012, Biogeosciences). However, since the initiation and climate history of each site are unknown, we assumed that all sites initiated 10100 years ago, with a constant present-day climate condition since their initiation and the peatland area hasn't changed, thus the simulated peat depth can't be compared to the measured depth.

The model was started with no catotelm layer, the carbon started to accumulate in the acrotelm layer, and as soon as carbon occurred in the acrotelm layer, a prescribed fraction of the acrotelm carbon was moved to the catotelm. When simulated water table (WT) drops below the acrotelm (WT_{min}), the whole acrotelm layer is supposed to decompose aerobically, as shown by Eq. 9, while the whole catotelm layer is still decomposing anaerobically. In the hydrology module, the total soil depth is 2m, we assumed that all layers in the peat soil profile hold peat-specific hydraulic properties, and there is no mineral soil beneath the peat soil. While the soil thermodynamics in the soil thermal module has 32 layers (38m), in which the top first 11 layers are identical to layers in hydrology, soil profiles in one grid cell are treated as mineral soil, and the dominant texture is used to define soil thermal properties.

Line 331-332: k_A and k_C are defined as fixed parameters, but line 319 says that they have a temperature dependence that is not shown in equations 5-8. These equations should show

the complete calculation, including temperature dependence etc.

We revised the equations on Page11, Line 318: ".....Decomposition of peat carbon is controlled by temperature (f_T) and parameterized as an exponential function:

 $f_T = Q_{10} exp((T-T_{ref})/10 \ \text{C})$ with $Q_{10} = 2.0$ and $T_{ref} = 30 \ \text{C}$ (Text S1). Soil carbon fluxes are given by:

$$\mathbf{F}_{AC} = \mathbf{k}_{p} \mathbf{f}_{T} \mathbf{C}_{A}$$
 (5)

$$R_{A,o} = \beta k_A f_T C_A , \qquad (6)$$

$$R_{A,a} = (1 - \beta) v k_A f_T C_A , \qquad (7)$$

$$R_{\rm C} = k_{\rm C} f_{\rm T} C_{\rm C} \tag{8}$$

Line 493-496: If this were the correct explanation, I would expect WT to be more accurately simulated in fens than in bogs. Was that the case?

We can't conclude that WT should be more accurately simulated in fens than bogs because we don't know the real amount of water input from non-peatland areas to peatland at fen sites. In this study, we routed all runoff from non-peatland soils into peatland. Considering that water table is relatively sensitive to the peatland area fraction in the grid cell (Fig. S11), it's hard to quantify if this water input setup caused greater errors in bogs than fens or not. The Taylor diagram (Fig. 3f) showed that there is no significant evidence for concluding that WT of fens are better simulated than bogs. We added a sentence on Page17, Line496 to point out the possible cause of this result: ".....an extra water source for bogs than only rainfall. However, the model did not perform better for fens (Fig. 3f), possibly because the amount of water that was routed into the fen was in error.".



Fig. S11. Sensitivity test of simulated water table to peatland area fraction in the grid cell, performed at the fen site FI-Lom.

Line 496-499: This seems like a very likely explanation to me, and something that could be tested by using a range of source-area/peatland-area ratios. Watershed analyses for the sites in question could provide some suggestions of realistic ratios.

We agree with the reviewer that watershed analyses could be helpful, but we feel that it's out the scope of this study. It could be considered for further developments of the model. Here, we performed a sensitivity test of simulated water table to peatland area fraction in the grid cell at one fen site (FI-Lom) to show the dependence of simulated water table on peatland area fraction (Fig. S11). We point out the dependence in the text, Page17, Line496: ".....3) WT depends on water input from surrounding non-peatland areas: the greater the peatland fraction in the grid cell, the smaller runoff input from other soils to the peatland, hence resulting in a deeper water table in the peatland (Fig. S11). ".

Line 515-516: This really highlights how the main peatland-related processes in the model are related to decomposition and respiration, not plant growth. Since that's the case, why is the evaluation so focused on photosynthesis? I think analysis of respiration fluxes would be much more informative, particularly in this case where WT would be expected to have an effect.

As NEE is the small residual of GPP and ER, wrong values of GPP could be one of major sources for errors in simulated NEE, especially when we only have one PFT to represents peatland vegetations. We agree with the reviewers that respiration fluxes are informative, thus we added analyses of ecosystem respiration. Please refer to our response to the first comment of the reviewer.

Line 531: Water use efficiency doesn't really fit with these other variables. It's a biological parameter, not a climate forcing variable like the other ones.

Water use efficiency and water balance were included because we would like to find a

variable / parameter that is possibly related to the optimized V_{cmax} , and may be used in the future to prescribe the spatial pattern of V_{cmax} in larger scale simulations in the future. So here we included not only climate forcing variables, but also these two biological parameters. We rephrased the name of Fig.S5 as: "Fig. S15. Relationship between optimized V_{cmax} and meteorological variables and biological parameters, as well as latitude of the sites location"

Line 536-537: It's surprising that there is no difference in Vcmax between fens and bogs, since those have very different vegetation types and productivities.

We recognize that bogs are precipitation-fed and nutrient-poor while fens are fed by precipitation and groundwater and can be either oligotrophic or eutrophic. However, previous studies have shown that along a bog-rich fen gradient in Alberta, Canada, the total above-ground net primary production exhibited a pattern of bog < poor fen < wooded moderate-rich fen> extreme rich fen> sedge fen (Szumigalski and Bayley, 1996, Wetlands), the productivity of the bog was not significantly lower than the poor fen and was even higher than the sedge and the extreme-rich fen. Also in Alberta, Thormann and Bayley (1997, Ecoscience) compared total aboveground plant production along a bog-fen-marsh gradient in Alberta, Canada, and found that the bog and the three fens (a lacustrine sedge fen, a riverine sedge fen and a floating sedge fen) had a similar NPP, the lacustrine sedge fen was even significantly less productive than the bog. The sites used in our study include wooded fens, wooded bogs, grassy fens and grassy bogs. Among them, we can't see a significant difference in dominant vegetation types between fens and bogs (we don't know relative abundances of grasses vs. shrubs vs. trees at each site though). We compared measured GPP of fens with that of bogs, there is no significant difference between them (P=0.63), as shown in the figure below.

Site	Туре	Aboveground biomass (kg/m2)	Dominant vegetation type		
DE-Bou	bog	grass dominated: 0.577;			
		heather and moss dominated: 0.517;	grasses, mosses		
		mixed: 0.303			
SE-Faj	bog	shrubs: 0.153; graminoids: 0.077;	ahruha araggag maggag		
		moses: 0.192	sinuos,grasses,mosses		
CA-Mer	bog	vascular: 0.356; mosses: 0.144	shrubs,mosses		
NO-And	bog		shrubs, grasses, mosses		
DE-Sfn	bog		trees,shrubs,grasses,mosses		
US-Bog	bog		trees, mosses		
SE-Deg	fen	vascular:0.049; mosses:0.065	shrubs,grasses,mosses		
CA-Wp3	fen	0.157	grasses,mosses		
CA-Wp2	fen	0.231	shrubs,grasses,mosses		
DK-Zaf	fen	0.471	grasses,mosses		
CZ-Wet	fen	0.57	grasses		
NO-Adv	fen	0.85	shrubs, grasses, mosses		
CA-Wp1	fen	1.08	trees,shrubs,mosses		
US-Los	fen	1.336	trees,shrubs,grasses		

DE-Spw	fen	trees
PL-Kpt	fen	grasses, reeds and ferns
DE-Zrk	fen	grasses
DK-NuF	fen	grasses,mosses
US-Fen	fen	grasses,forbs
FI-Sii	fen	shrubs,grasses,mosses
FI-Lom	fen	shrubs,grasses,mosses



Line 540-541: This really seems like it could be compensating for some other error related to vegetation biomass, LAI, or productivity. I would expect higher biomass and LAI in warmer areas, which would drive exactly this type of relationship. I think this should be investigated since the optimization of Vcmax could be masking other important model issues.

The measured LAI indeed is larger in warmer areas, but we would like to mention that there is no systematic bias in LAI or biomass, as shown in Fig.S1. Verheije et al. (2013, Biogeosciences) demonstrated that Earth system models could be improved by taking plant traits variations within PFTs into account, and proposed relationships between trait parameters and the climate, which can be used to define the parameter values for each grid cell. Considering that there is no available observational-based trait-climate relationships that can be used for peatland vegetations, we optimized V_{cmax} at each site and built the relationship between the optimized V_{cmax} and the latitude (temperature), which showed better performance than using a mean value. The peat PFT in our study represents an average of the ecosystem, not a specific plant type. A broad decrease of V_{cmax} with latitude in the northern hemisphere has also been documented by Walker et al. (2017, New Phytologist), assuming that V_{cmax} was constrained by the rate of N uptake, with the rate of N uptake calculated as a function of soil C, N and mean annual air temperature. We note this in the text on Page20, Line587: ".....relationship with the latitude of chosen peatland sites location. A decrease of V_{cmax} with latitude in the northern hemisphere, like the one inferred from optimized sites values, has also been documented by

Walker et al. (2017), who assumed that V_{cmax} was constrained by the rate of N uptake, with the rate of N uptake calculated as a function of soil C, N and mean annual air temperature. We speculate the dependence of optimized V_{cmax} on latitude found in Sect. 4.2 can be attributed to.....".

Line 549: Why not use this mean value of 40 in the previous comparison, instead of the default value of 16?

The mean value of 40μ mol m⁻² s⁻¹ is used in the revised manuscript. Please refer to our responses to the second comment.

Line 560-561: Why are only these two sites discussed and shown in the figure? Was the relevant data not available for other sites, or are these just being used as illustrative examples? We have data for other sites. The underestimation of soil temperature in winter and overestimation in summer occurred at most of these sites. DK-Nuf and CA-Wp1 are just used as illustrative examples. We corrected the text on Page19, Line560: ".....soil temperature was underestimated in winter and overestimated in summer by our model (Fig. 7 and 8, results from sites DK-Nuf and CA-Wp1 are shown as illustrative examples)."

Line 564-566: The suggestion that the issues are due to errors in snow density implies that the snow mass was correct in the model. Is that true?

We didn't validate the simulated snow mass because of lack of available data. We rephrased the text on Page19, Line564: ".....can be caused by the bias in snow processes of the model, such as underestimation of snow mass, and/or overestimation of snow density and...".

Line 582-585: Even if optimized Vcmax is an average for the ecosystem rather than a species-specific value, it should be comparable with the observed range among different species that exist in these systems. Other peatland models should definitely be comparable, because any peatland model would be representing an average plant type. I don't think this is a satisfying explanation for not comparing the optimized estimates with measurements. It's just as likely that the model underestimates LAI and needed to tune Vcmax higher to compensate. I don't find any of the three explaination below particularly convincing, and I think bias in LAI or plant biomass is a likely explanation that should be tested.

The reviewer is right, the optimized V_{cmax} should be compared with the observed range among different species. Therefore we added these sentences on Page20, Line582: ".....The V_{cmax} values estimated in this study ranged from 19 to 89 µmol m⁻² s⁻¹, with a mean value of 40 µmol m⁻² s⁻¹. These values were not fully comparable with values reported for a specific vegetation type, as they are averages for all plants growing in the peatland ecosystem. As stated in Sect. 2.2, observed V_{cmax} varies strongly among different species and sites. V_{cmax} of mosses at the Old Black Spruce site (Canada) varied from 5 to 14 µmol m⁻² s⁻¹ (Williams and Flanagan, 1998), In a nutrient addition experiments conducted by Bubier et al. (2011), V_{cmax} for ericaceous shrubs in a temperate bog ranged from 67 to 137µmol m⁻² s⁻¹, with V_{cmax} for *Vaccinium myrtilloides*, *Ledum groenlandicum* and *Chamaedaphne calyculata* valued at 84.6 ± 13.5 µmol m⁻² s⁻¹, 78.1 ± 13.4 µmol m⁻² s⁻¹, and 132.1 ± 31.2 µmol m⁻² s⁻¹ in the plots with no nutrient addition. The optimized model V_{cmax} in our study was within the range of these observations. Meanwhile, the values we inferred from sites to match peak GPP are comparable to those used in other land surface models: the McGill wetland model used a value of 17 µmol m⁻² s⁻¹ for evergreen shrubs (St-Hilaire et al., 2010); the CLASS-CTEM model (Wu et al., 2016) used 60, 50, 40 µmol m⁻² s⁻¹ for evergreen shrubs, deciduous shrubs and sedges, respectively; the values for mosses in these two models were adapted from the study of Williams and Flanagan (1998). ".

Line 591-592: Does ORCHIDEE not already take the influence of temperature on photosynthesis into account?

ORCHIDEE does take the influence of temperature on photosynthesis into account by parameterizing the temperature dependences of Michaelis-Menten constants, CO₂ compensation point following Medlyn et al. (2002, Plant, cell & environment). And temperature acclimation of photosynthesis rates constants is included in ORCHIDEE following Yin et al. (2009, NJAS-Wageningen J. Life Sci.). We thus removed the following sentences on Page20, line 591-592: ".....2) with an adequate water supply, leaves open their stomata in response to warm environments, leading to a higher photosynthetic efficiency (Chapin III et al., 2011);".

Line 593: If the issue were nutrient availability, I would expect strong contrasts in Vcmax between fen and bog ecosystems, which did not appear to be the case in this study.

As we mentioned above, the sites used in this study include wooded fens, wooded bogs, grassy fens and grassy bogs, among them, there is no significant difference in dominant vegetation types between fens and bogs. Meanwhile, there is neither significant difference in measured biomass between fens and bogs (P=0.097) nor significant difference in measured GPP (P=0.63).

Line 603-632: This is a nice review of observed drought effects on peatlands, but the paper doesn't demonstrate whether the model can reproduce any of these effects. Such a demonstration would be very informative.

We added these sentences to demonstrate results of the model on Page22, Line628: ".....and growth of peatland vegetation was not constrained by water table depth in the model. Therefore, the sensitivity of GPP to WT fluctuations in observations was not included in the model. As a consequence, the model neither captured the reported decrease of photosynthesis due to drought at CA-Wp3 (Adkinson et al., 2011) and SE-Faj (Lund et al., 2012), nor the increase of photosynthesis as a result of lower water table at CA-Wp1 (Flanagan and Syed, 2011). However, the model can reproduce the pattern that above a critical level (acrotelm depth), peat respiration decreases with increasing WT (Fig.5, Fig.S13), as reported at site CA-Mer and US-Los (Lafleur et al., 2005; Sulman et al., 2009). ".

Line 630-632: It would be better to show that the model reproduces this pattern (in a figure) rather than just asserting that it can.

The decrease of soil respiration with increasing WT (shallower) was shown in Fig.5 and Fig. S13. We added this sentence on Page22, Line630: "..... The model can reproduce the

pattern that above a critical level (acrotelm depth), peat respiration decreases with increasing WT (Fig.5, Fig.S13), as reported at site CA-Mer and US-Los (Lafleur et al., 2005; Sulman et al., 2009).".

Line 634-635: If GPP was captured well but NEE was not, then the difference must be due to simulated respiration. This is another case where more analysis of simulated respiration would be very helpful.

Ecosystem respiration was relatively well captured by the model. We added these sentences on Page22, Line634: ".....variations in GPP (with $r^2 = 0.75$, 0.86, and 0.93, respectively) and ER (with $r^2=0.78$, 0.86, and 0.89, respectively), but were less able to reproduce variations in NEE (with $r^2 = 0.38$, 0.61, and 0.27, respectively). Note that in the two-layer soil carbon scheme, the dependence of soil respiration on temperature was parameterized as an exponential function of the soil layers-weighted average temperature (Text S1).and values of Q₁₀ coefficient depend on the soil depth (Lafleur et al., 2005; D'Angelo et al., 2016). Small-scale peatland surface heterogeneities are not included in the model,"

Line 666: This implies that water table is not an important feature of carbon cycling according to this model. This seems very inconsistent with the observational literature showing that peatland CO2 fluxes are quite responsive to water table fluctuations (much of which is cited in this manuscript). Some papers have demonstrated that compensating responses of GPP and respiration (e.g. both increasing under a drying trend) can cause NEE to be insensitive to water table fluctuations (e.g. Sulman et al. 2010), but the paper doesn't really demonstrate that the model is reproducing those compensating responses. Given the centrality of water table and hydrology in our understanding of peatland carbon cycling, I think this conclusion that water table isn't actually that important needs to be investigated in more detail, especially in how it affects peat decomposition and ecosystem respiration in the model.

The point we were trying to make here is that although water table was poorly simulated by the model, it was good enough to simulate ER (NEE) properly. With water table being forced to be equal to observed values in S2, there were no large improvements in simulated ER, NEE (Table5, Table6, Fig. S13). This is because the oxic decomposition in the acrotelm (β), which is the main component of soil respiration, was calculated by comparing the height of the acrotelm with the WT depth, though absolute values of water table depth in S1 and S2 were quite different (Fig. S8), β were not so different. We took Lompolojänkkä fen site (FI-Lom) as an example, in which WT was most severely underestimated. As shown by Fig. S12, difference between β of S1 and S2 only occurred during short periods and mainly in winter when decompositions were inhibited by the low temperature. We performed an additional simulation (S3), in which we assumed that water table was more severely underestimated by the model (water table used in S3 was consistently 20cm deeper than in S1), thus the acrotelm was more exposed to the air in S3 (Fig. S12). S3 showed much larger ecosystem respiration and hence smaller carbon sequestration than S1. We clarified this by added these sentences on Page18, Line524: ".....an overestimation (more negative values) of NEE in the warm period (May-September). The influence of WT on respiration was parameterized as the separation of oxic (β in Eq. 6) vs. anoxic (1- β in Eq. 7) decomposition in the acrotelm. Although absolute values of simulated WT in S1 and WT_{obs} in S2 were quite different (Fig. S8), the values of β were not very different (Fig.S12). Therefore the simulated WT was good enough to properly replicate ER (Fig.S13). An additional simulation (S3) performed at FI-Lom showed that if WT was more severely underestimated, e.g. WT in S3 was consistently 20 cm deeper than in S1, the acrotelm was exposed to oxygen for longer time, resulting in larger ER and hence smaller carbon sequestration in S3 (Fig.S12, Fig.S13).". We rephrased the sentences in abstract on Page3, Line105: ".....likely due to the uncertain water input to the peat from surrounding areas. However, the poor performance of WT did not greatly affect predictions of ER and NEE.", and the sentences in conclusion on Page23, Line665: ".....instead of calculated by the model, was small, indicating that the simulated WT was reliable to predict ER and NEE properly."



Fig. S12. The fraction of the acrotelm where carbon decomposes under oxic conditions (β) at Lompoloj änkk ä fen site (FI-Lom). S1: simulated water table (WT) were used in the carbon module; S2: observed water table (WT_{obs}) were used in the carbon module; S3: assumed that water table were 20cm deeper than simulated results, thus (WT–20cm) were used in the carbon module.



Fig. S13. Cumulative ER (left figure) and NEE (right figure) at Lompolojänkkä fen site (FI-Lom). S1: simulated water table (WT) were used in the carbon module; S2: observed water table (WT_{obs}) were used in the carbon module; S3: assumed that water table were 20cm deeper than simulated results, thus (WT–20cm) were used in the carbon module.

Line 670-671: The paper definitely did not establish that nitrogen availability was the explanation for the latitudinal dependence. It was one of several proposed explanations. In fact, I think it's unlikely to be the explanation because it did not vary consistently with fen/bog type, which is closely related to nitrogen availability.

Not all fens in this study are nutrient rich, for example, SE-Deg (Peichl et al., 2014, Environ. Res. Lett.), FI-Sii (Aurela et al., 2007, Tellus), CA-Wp2 (Adkinson et al., 2011, J. Geophys. Res. Biogeosciences) are oligotrophic fens, thus there is a large variation in V_{cmax} of fens. And there is no significant difference in biomass, GPP between fens and bogs. Meanwhile, Walker et al. (2017, New Phytologist) found that V_{cmax} values decreased with latitude in the northern hemisphere if the rate of nitrogen uptake was parameterized as a function of soil C, N, and mean annual air temperature. Thus, we can't rule out the possibility that the relationship was caused by nitrogen availability.

Table 2: In addition to bog/fen type, it would be informative to include something about the dominant vegetation type (grass, shrub, forested) and maybe aboveground biomass or LAI if available

We included the dominant vegetation type and LAI, and aboveground biomass in the Table2, detailed description of the sites can be found in the supplement material.