

# 1 **Sensitivity Analysis of a Coupled Hydrodynamic-Vegetation Model** 2 **Using the Effectively Subsampled Quadratures Method (ESQM v5.2)**

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5 Tarandeep S. Kalra<sup>1</sup>, Alfredo Aretxabaleta<sup>2</sup>, Pranay Seshadri<sup>3</sup>, Neil K. Ganju<sup>2</sup>, and Alexis Beudin<sup>4</sup>

6 <sup>1</sup>Integrated Statistics, contracted to the U.S. Geological Survey, Woods Hole, MA 02543, USA

7 <sup>2</sup>U.S. Geological Survey, Woods Hole, MA 02543, USA

8 <sup>3</sup>University of Cambridge, Cambridge, England, CB2 1PZ, United Kingdom

9 <sup>4</sup>University of Bordeaux, Bordeaux, France

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11 *Correspondence to:* Tarandeep S. Kalra (tkalra@usgs.gov)

12 **Abstract.** Coastal hydrodynamics can be greatly affected by the presence of submerged aquatic vegetation. The effect of  
13 vegetation has been incorporated into the Coupled-Ocean-Atmosphere-Wave-Sediment Transport (COAWST) Modeling  
14 System. The vegetation implementation includes the plant-induced three-dimensional drag, in-canopy wave-induced  
15 streaming, and the production of turbulent kinetic energy by the presence of vegetation. In this study, we evaluate the sensitivity  
16 of the flow and wave dynamics to vegetation parameters using Sobol' indices and a least squares polynomial approach referred  
17 to as Effective Quadratures method. This method reduces the number of simulations needed for evaluating Sobol' indices and  
18 provides a robust, practical, and efficient approach for the parameter sensitivity analysis. The evaluation of Sobol' indices  
19 shows that kinetic energy, turbulent kinetic energy, and water level changes are affected by plant stem density, height, and to  
20 a lesser degree, diameter. Wave dissipation is mostly dependent on the variation in plant stem density. Performing sensitivity  
21 analyses for the vegetation module in COAWST provides guidance to optimize efforts and reduce exploration of parameter  
22 space for future observational and modeling work.

## 23 **1 Introduction**

24 The presence of aquatic vegetation (e.g., mangroves, salt marshes, and seagrass meadows) provides several ecological benefits  
25 including nutrient cycling, habitat provision, and sediment stabilization (Costanza et al., 1997). Vegetation provides habitat  
26 for many species of epiphytes, invertebrates, and larval and adult fish (Heck et al., 2003). Seagrass meadows reduce sediment  
27 resuspension, thereby stabilizing bottom sediment, increasing light penetration, and improving water clarity in a positive  
28 feedback loop (Carr et al., 2010). In addition, aquatic vegetation provides coastal protection by absorbing wave energy  
29 (Wamsley et al., 2010).

30 One approach to implement the influence of aquatic vegetation on circulation is by increasing the bottom roughness  
31 coefficient (Ree, 1949; Morin et al., 2000). More recently, the effect of vegetation being studied through parameterization as  
32 form drag as opposed to skin friction in 2-D depth averaged models (Chen et al., 2007; Le Bouteiller and Venditti, 2015). To  
33 account for 3-D vertical structures, estuary-scale models have implemented both mean and turbulent flow impacts of vegetation  
34 (Temmerman et al., 2005; Kombiadou et al. 2014; Lapetina and Sheng, 2015). In addition to impacts on the flow-field, the  
35 presence of vegetation also results in wave attenuation. The decay of wave height over vegetation has been simulated by  
36 enhancing bed roughness (Möller et al., 1999; de Vriend, 2006; Chen et al., 2007). A more physical description of wave  
37 attenuation due to vegetation was developed by Dalrymple et. al (1984), who approximated wave energy loss due to stalks  
38 approximated as cylinders. This approach has been applied in spectral wave models and calibrated against flume experiment  
39 results (Mendez and Lozada, 2004; Suzuki et al., 2012; Wu, 2014; Bacchi et al., 2014).

40 Recently, Beudin et al. (2017) implemented the effects of vegetation in a vertically varying water column through  
41 momentum extraction and turbulence dissipation and generation using a 3D hydrodynamic model and accounting for wave  
42 dissipation due to vegetation in a spectral wave model. The modeling approach was implemented and tested within the open  
43 source COAWST (Coupled-Ocean-Atmospheric-Wave-Sediment Transport) modeling system that couples hydrodynamic and  
44 wave models (Warner et al, 2010). The vegetation module was based on modifications to the flow field resulting from three-  
45 dimensional drag, in-canopy wave-induced streaming, and production of turbulent kinetic energy in the hydrodynamics model  
46 (ROMS), along with energy dissipation and resultant hydrodynamic feedback from the wave model (SWAN).

47 The vegetation module requires the user to input a given set of plant properties (stem density, height, diameter, and  
48 thickness). These vegetation properties can be highly variable depending on the season and environment, yet obtaining a full  
49 set of measurements in realistic settings is impractical. Identifying which properties have the greatest influence on the resulting  
50 flow dynamics can reduce the amount of observational data required to robustly parameterize the model, and/or reduce the  
51 number of runs required in a model ensemble to quantify the uncertainty associated with data gaps. Our study aims to perform  
52 a systematic sensitivity analysis to quantify the effect of changing the vegetation properties on the resulting hydrodynamic  
53 output. The results of the sensitivity analysis can be used to select and rank the most important parameters for calibration. Two  
54 conditions are required for the model to display a significant sensitivity: (1) a sufficient modification of one of the forcing  
55 parameters and (2) a change in the leading terms of the dynamic equations of the model. While modifying the forcing  
56 parameters by a sufficient amount is required, the modification should remain within the natural range of variability of the  
57 parameters.

58 Several mathematical techniques have been utilized to perform sensitivity analysis. Bryant (1987) applied scaling  
59 analysis to an idealized domain and forcing and found that closure parameters such as vertical diffusivity and wind stress curl  
60 were important controlling factors in thermohaline circulation. Bastidas et al. (1999) used multicriteria methods to find the  
61 sensitivity of land surface scheme models that couple biosphere-atmosphere interactions. The input variables (such as  
62 precipitation, air temperature and humidity etc.) predict the evolution of soil skin temperature, soil moisture, etc. The input  
63 parameters obtained from the sensitivity analysis of the model showed consistency with physical properties for two different

64 field sites and helped identifying insensitive parameters that led to an improvement in model description. Fennel et al. (2001)  
65 incorporated adjoint methods to perform sensitivity studies to refine ecological model parameters such that the underlining  
66 model can be applied to a wider range of conditions. Mourre et al. (2008) performed multiple simulations based on realistic  
67 variation of a forcing field to calculate the influence of model parameters on sea surface salinity. The metric used to measure  
68 the sensitivity was based on RMS difference between the reference and modified model parameter. The results showed that  
69 lateral salt diffusivity had the strongest impact on surface salinity model response. Rosero et al. (2009) investigated the  
70 sensitivity of three different versions of a land satellite model (Noah LSM) applied to nine different sites based on different  
71 conditions (soil, vegetation, and climate). They utilized the Monte Carlo method to generate the first order Sobol' indices  
72 (Sobol', 1993), to estimate the model sensitivity. The results showed that the optimal parameter values differed varied between  
73 different versions of the models and for different sites along with quantifying the nature of interactions between parameters.  
74 One of the challenges associated with a Monte Carlo approach to computing the Sobol' indices is the large number of model  
75 evaluations required for approximating conditional variance.

76 All these studies highlight various approaches to perform sensitivity analysis. Saltelli et al. (2008) provided a  
77 comparison of different sensitivity analysis methodologies and the optimal setup for specific combinations of parameters and  
78 model. Ultimately, the choice of sensitivity analysis methodology depends on multiple factors such as the computational cost  
79 of running the model, the characteristics of the model (e.g., nonlinearity), the number of input parameters, and/or the potential  
80 interactions between parameters. Saltelli et al. (2008) described variance-based techniques as providing the most complete and  
81 general pattern of sensitivity for models with a limited number of parameters, such as the vegetation module in COAWST.  
82 Sobol' indices, as a form of variance-based sensitivity analysis, provide a decomposition of the variance of a model into  
83 fractions that can be assigned to inputs or combinations of inputs. However, techniques involving the estimation of Sobol'  
84 indices (such as Monte Carlo methods) are expensive.

85 To reduce the computational cost and have desirable accuracy, techniques that involve approximating the global  
86 response of the model with a polynomial, and then using its coefficients to estimate the Sobol' indices can be utilized (Sudret,  
87 2008). In this paper, we use a set of least squares polynomial tools based on subsampling to estimate our global polynomial  
88 response (Seshadri et al., 2017a). Then, the coefficients of the polynomial are used to compute the Sobol' indices. These tools  
89 are implemented in the open source package Effective Quadratures Method (EQ) (Seshadri et al., 2017b), and our current work  
90 provides one of the first applications of this methodology to quantify sensitivity of input parameters in coastal models.  
91 Therefore, the goal of the present work is to take advantage of the EQ method to provide Sobol' indices that quantify the  
92 sensitivity of the flow and wave dynamics to vegetation parameters in COAWST model. The paper is organized as follows:  
93 the methods are discussed in section 2, including the numerical model with vegetation model (COAWST), Effective  
94 Quadratures method to estimate Sobol' indices, and simulation design; in section 3 we present the results of sensitivity analysis  
95 from various simulations; in section 4 we discuss the impact of these results; and finally in section 5 we summarize our work  
96 and outline areas of future research.

97

## 98 **2 Methods**

### 99 **2.1 COAWST implementation of vegetation model parameterization**

100 Beudin et al. (2017) implemented a hydrodynamic-vegetation routine within the open-source COAWST numerical  
101 modeling system. The COAWST framework utilizes ROMS (Regional Ocean Modeling System) for hydrodynamics and  
102 SWAN (Simulating WAVes Nearshore) for modeling waves coupled via the Model Coupling Toolkit (MCT) (Warner et al.,  
103 2008).

104 ROMS (Regional Ocean Modeling System) is a three-dimensional, free surface, finite-difference, terrain-following  
105 model that solves the Reynolds-Averaged Navier-Stokes equations using the hydrostatic and Boussinesq assumptions  
106 (Haidvogel et al., 2008). The transport of turbulent kinetic energy and generic length scale are computed with a generic length  
107 scale (GLS) two-equation turbulence model. SWAN (Simulating WAVes Nearshore) is a third-generation spectral wave model  
108 based on the action balance equation (Booij et al., 1999). The effect of submerged aquatic vegetation in ROMS is to extract  
109 momentum, add wave-induced streaming, and generate turbulence dissipation. Similarly, the wave dissipation due to  
110 vegetation modifies the source term of the action balance equation in SWAN. Sub-grid scale parameterizations account for  
111 changes due to vegetation in the water column extending from the bottom layer to the height of the vegetation in the flow  
112 model, while SWAN accounts for wave dissipation due to vegetation at the seafloor. The parameterization of SWAN to account  
113 for wave dissipation implemented by Suzuki et al. 2012 has the same effect as energy dissipation.

114 The parameterizations used to implement the effect of vegetation in both ROMS and SWAN models are mentioned  
115 in Table 1 and detailed in Beudin et al. (2017). The coupling between the two models occurs with an exchange of water level  
116 and depth averaged velocities from ROMS to SWAN and wave fields from SWAN to ROMS after a fixed number of time  
117 steps (Fig. 1). The vegetation properties are separately input in the two models at the beginning of the simulations.

118

### 119 **2.2 Method for Sensitivity Analysis: Polynomial Least Squares**

120 Polynomial techniques are ubiquitous in the field of uncertainty quantification and model approximation. They estimate the  
121 response of some quantity of interest with respect to various input parameters using a global polynomial. From the coefficients  
122 of the polynomial the mean, variance, skewness, and higher order statistical moments can be calculated (see Smith 2014;  
123 Geraci et al. 2016). In this paper, our interest lies in statistical sensitivity metrics called first order Sobol' indices (Sobol',  
124 1993) that are derived from the conditional variances of the parameters of the model. These indices are the same in number as  
125 the input parameters to the model and quantitatively rank the input parameters based on their contribution to the resultant  
126 model output. Thus, model output is more sensitive to parameters that exhibit higher first order Sobol' index value. Second  
127 order and third order Sobol' indices may also be computed. The sum of the first, second, and third order Sobol' indices should  
128 sum to unity, therefore if the first order indices are themselves close to unity, it indicates the higher order interaction between  
129 model input parameters is weak.

130 In this paper, the first order indices are computed from a global polynomial model using Effectively Subsampled  
131 Quadratures Method (ESQM v5.2; Seshadri et al., 2017b). There are two attributes to any data-driven polynomial model: the  
132 choice of the polynomial basis and the strategy for estimating the coefficients of the polynomial. The basis used in Effective  
133 Quadratures are orthogonal polynomials; i.e. orthogonal with respect to the weight of the input parameter. For example, if one  
134 of the input parameters is prescribed with a Gaussian distribution, then a Hermite orthogonal polynomial basis would be used;  
135 likewise for a uniform distribution, Legendre polynomials are used. In the current work, a uniform distribution is assumed for  
136 the input parameter space. The rationale behind selecting polynomials that are orthogonal with respect to the input weight is  
137 that it reduces the number of model evaluations required for estimating statistical moments. Details on the exponential  
138 convergence in moments when matching the orthogonal polynomial with its corresponding weight can be found in Xiu et al.  
139 (2002).

140 The coefficients for the polynomial expansion are typically approximated using an integral over the input parameter  
141 space using quadrature rules. When the number of input parameters is greater than one, tensor grid or sparse grid based  
142 quadrature rules may be used to approximate these integrals. However, the cost of tensor grids grows exponentially with  
143 dimension, i.e., a four-point quadrature rule in three dimensions has  $5^3$  points, in four dimensions has  $5^4$  points and so on.  
144 While some alleviation can be obtained using sparse grids, in this paper a more efficient sampling technique is used: effectively  
145 subsampled quadratures (abbreviated to Effective Quadratures).

146 The method of Effective Quadratures determines points for approximating the integral by subsampling well-chosen  
147 points from a tensor grid, and evaluating the model at those subsamples. These well-chosen points are obtained via a QR  
148 column pivoting heuristic (Seshadri et al. 2017a). Once the coefficients are estimated, the Sobol' indices can be readily  
149 computed (Sudret 2008).

150

### 151 **2.3 Range of input vegetation properties for sensitivity analysis**

152 Prior to performing the simulations for estimating Sobol' indices described above, a range of vegetation inputs that would  
153 impact the model response needs to be chosen. Kennish et al. (2013) constrained annual variation of three of the four vegetation  
154 properties (stem density, height, diameter) based on *Zostera marina* growth in the Barnegat Bay-Little Egg Harbor estuary.  
155 The thickness of *Z. Marina* is selected to be an order of magnitude lesser than its diameter based on Lakrum et al., 2007. Based  
156 on this, the range of the vegetation model inputs evaluated is as follows:

- 157 1. Stem density ( $n_v$ ) = [38.2 - 250.4] stems/m<sup>2</sup>.
- 158 2. Height ( $l_v$ ) = [0.16 - 0.32] m.
- 159 3. Diameter ( $b_v$ ) = [1.0 - 10.0] mm.
- 160 4. Thickness ( $t_v$ ) = [0.1 - 1.0] mm.

161 For the sensitivity analysis, a combination of these ranges of inputs (Table 2) is chosen to configure different simulations in  
162 an idealized test case, described below. In addition to these four vegetation properties, the vegetative model requires an input  
163 of drag coefficient ( $C_D$ ) in the flow model and the wave model. However, variations on  $C_D$  are unlikely to be measured in the

164 field and thus users could rely on the published literature for an appropriate choice based on the type, shape, and flexibility of  
165 the vegetation under study.

166

#### 167 **2.4 Test case configuration**

168 An idealized rectangular model domain of 10 km by 10 km with a 3 m deep basin is chosen. The grid is 100 by 100 in the  
169 horizontal (100 m resolution) and has 60 vertical sigma-layers (uniformly distributed) leading to 0.05 m resolution in the  
170 vertical. The vertical resolution of 0.05 m allows a plant height of 0.27 m to be distributed over 6 vertical layers while the  
171 shortest height is restricted to 2 vertical layers. A square patch of vegetation (1 km by 1 km) is placed in the middle of the  
172 domain. The ROMS barotropic and baroclinic time steps are respectively *0.05 s* and *1 s*, while the SWAN time step and the  
173 coupling interval between ROMS and SWAN are 10 min. The friction exerted on the flow by the bed is calculated using the  
174 SSW bbl in the bottom boundary layer formulation (Warner et al., 2008). The bottom boundary layer roughness is increased  
175 by the presence of waves that produce enhanced drag on the mean flow (Madsen, 1994; Ganju and Sherwood, 2010). The  
176 vegetative drag coefficients ( $C_D$ ) in the flow model and the wave model are set to 1 (typical value for a cylinder at high  
177 Reynolds number). The bed roughness is set to  $z_0=0.05$  mm, which corresponds to a mixture of silt and sand (Soulsby, 1997).  
178 The turbulence model selected is the  $k - \epsilon$  scheme (Rodi, 1984).

179 The model is forced by oscillating the water level on the northern edge with a tidal amplitude of 0.5 m and a period  
180 of 12 h. Waves are also imposed on the northern edge with a height of 0.5 m, directed to the south (zero angle), with a period  
181 of 2 s. The test case setup is similar to the one used by Beudin et al., (2017). The test case setup is simulated for 2 days to  
182 obtain a tidally steady state solution. These simulations require 40 CPU hours on Intel Xeon(R) X5650 2.67 GHz processors  
183 running on 24 parallel processors.

184

#### 185 **2.5 Choice of COAWST model response to vegetation inputs**

186 The output parameters used to investigate the vegetation model sensitivity are chosen to reflect the first order effects of  
187 vegetation on the hydrodynamics and waves. The results for the model response are computed for the last tidal cycle (a total  
188 of 3 tidal cycles are required for achieving steady state). The presence of vegetation affects the output parameters in different  
189 physical ways (Table 1) as:

- 190 1. Wave energy dissipation: Vegetation dissipates wave energy by reducing wave height, increasing wavelength, and reducing  
191 wave steepness.
- 192 2. Kinetic energy: A drag force is generated by plants. This leads to a decrease in kinetic energy within and behind the  
193 vegetation patch.
- 194 3. Water level: As the wave energy (and momentum) flux decreases due to bottom friction, the mean water level increases to  
195 balance the decrease in wave and kinetic energy. The flow decelerates in front of the patch and in the wake of the patch while  
196 it accelerates around the edges of the patch, leading to a water level gradient.

197 4. Turbulent kinetic energy (TKE): Reduced turbulent kinetic energy in front and within the patch. The enhancement of TKE  
198 inside the boundary layer is not captured with the current resolution.

199 The response impact to change in the inputs during each simulation is computed by calculating the percentage  
200 difference of model response for each simulation from the minimum value of all the simulations. The model response is  
201 obtained in and around the vegetation patch and averaged over the last tidal cycle. The change in model response of water  
202 level is computed by finding the maximum water level difference in and around the vegetation patch. In addition, the variability  
203 of model response with given vegetation inputs in different simulations is calculated through standard deviation of model  
204 response in and around the vegetation patch over the last tidal cycle. The standard deviation in TKE is depth averaged to  
205 provide a 2-D field.

206

## 207 **3 Results**

### 208 **3.1 Setting up simulations with different vegetation inputs**

209 Using the range of input parameters described above (section 2.3), and assuming all the inputs are uniformly distributed over  
210 their ranges, a matrix of design of experiment values (Table 2) was determined using Effective Quadratures. A total of 15  
211 simulations were found to be required, corresponding to the number of coefficients in a 4D polynomial with a maximum order  
212 of 2. In general, the number of coefficients  $n$ , is given by the formula

$$213 \quad n = \binom{d + o}{d},$$

214 where  $d$  is the number of dimensions and  $o$  is the maximum order (assuming it is isotropic across all dimensions).

215

### 216 **3.2 Model response from all the simulations**

217 The 15 simulations (parameter choice of each simulation in Table 2) are performed to provide model response from the four  
218 chosen output variables.

219 1. Wave dissipation: The percentage change in wave dissipation from all simulations relative to the minimum value of wave  
220 dissipation varied between 75 and 600% (Fig. 3(a)). Simulation 2, which includes the combination of smallest stem density  
221 ( $n_v = 62.1$  stems/m<sup>2</sup>) and shortest height (0.174m), incurred the least wave dissipation, while simulation 3, which involved the  
222 combination of greatest stem density ( $n_v = 226.5$  stems/m<sup>2</sup>) and tallest height (0.295 m), resulted in greatest wave dissipation.  
223 The greatest amount of variability in wave dissipation occurs in front of the vegetation patch (Fig. 4), where the greatest  
224 amount of wave energy is dissipated due to the presence of the vegetation patch.

225 2. Kinetic energy: The percentage change in kinetic energy from all simulations relative to the minimum kinetic energy varied  
226 between 5.0 and 34.0 % (Fig. 3b). Simulation 8, performed with an intermediate value of stem density along with the largest  
227 values of height, diameter, and thickness, results in the least amount of kinetic energy (lowest velocities). Simulation 2 causes  
228 the least amount of extraction of momentum with a combination of smallest plant stem density, height, and diameter values,  
229 resulting in the greatest kinetic energy in and around the vegetation patch. The variability in kinetic energy from all the

230 simulations is observed at a cross section along the vegetation patch (Fig. 5). Variability is observed throughout the water  
231 column where the vegetation patch exists. Similar to the variability in wave dissipation, the greatest amount of variability  
232 occurs in front of the vegetation patch. The region of maximum variability occurs at a distance of (0.25 – 1.25) m above bed.  
233 3. Water level: The percentage change in maximum water level difference from all simulations relative to the minimum of the  
234 maximum water level difference varied between 3.0 and 18.0 % (Fig. 3c). The minimum water level gradient is obtained from  
235 simulation 13 that includes a combination of plant inputs ( $n_v = 62.1$  stems/m<sup>2</sup> (minimum),  $l_v = 0.174$  m (intermediate),  
236 diameter=6 mm (minimum), and  $t_v = 0.3$  mm (minimum). Simulations 2, 6, and 11 also give relatively low values of water  
237 level gradient. Simulations 2 and 6 both involve the smallest plant stem density values (i.e.  $n_v = 62.1$  stems/m<sup>2</sup>). On the other  
238 hand, simulation 8 ( $n_v = 144.3$  stems/m<sup>2</sup>,  $l_v = 0.3$  m (maximum),  $b_v = 9$  mm, and  $t_v = 0.9$  mm) accounts for the greatest value of  
239 maximum water level difference. The variability in water level (Fig. 6) is highest around the lobes of the vegetation patch  
240 where the water level adjusts owing to changes in velocity around the patch. Behind the vegetation patch, the variability  
241 increases as the water level adjusts to the change in flow conditions.

242 4. Turbulent Kinetic Energy (TKE): The percentage change in TKE from all simulations relative to the minimum TKE varied  
243 between 0.5 and 10.0 % (Fig. 3d). Simulation 8 (combination of plant  $n_v = 144.3$  stems/m<sup>2</sup>,  $l_v = 0.3$  m,  $b_v = 9$  mm and  $t_v = 0.9$   
244 mm) gives the greatest amount of TKE, similar to what was observed with kinetic energy change. The smallest value of TKE  
245 is obtained in and around the vegetation patch in simulation 2 with a combination of plant inputs ( $n_v = 62.1$  stems/m<sup>2</sup>,  $l_v = 0.2$   
246 m,  $b_v = 3$  mm, and  $t_v = 0.6$  mm). Simulation 2 causes the least amount of dissipation of turbulence with a combination of  
247 smallest plant stem density, height and diameter. The variability in TKE peaks occurs in front of the vegetation patch (Fig. 7)  
248 where the different simulations dissipate turbulence to substantially different degrees. The changes in turbulence mixing caused  
249 by the presence of the vegetation patch are close to zero inside the patch (all simulations dissipate similar amounts of  
250 turbulence).

251

### 252 3.3 Quantifying sensitivity using Sobol' Indices

253 Following the variability in model response from different simulations, the sensitivity to input vegetation parameters can be  
254 quantified with the use of first order Sobol' indices that are obtained by taking advantage of the Effective Quadratures  
255 approach. Sobol' indices are individually computed for all the model responses. The first order Sobol' indices for all the model  
256 responses (Table 3) add up to more than 0.9. This results indicates they account for 90% of the variability in model response  
257 for the given vegetation property inputs, and the variability captured by second order and third order indices is relatively low.  
258 The model is most sensitive (Table 3) to plant stem density ( $n_v$ ) and height ( $l_v$ ) over the range of parameters considered; these  
259 two inputs account for over 80% of the sensitivity to all model outputs. The vegetation diameter ( $b_v$ ) accounts for 12-15 % of  
260 model sensitivity to kinetic energy, water level change, and turbulent kinetic energy. Thickness ( $t_v$ ) showed the least impact  
261 on all the chosen model outputs.

262

263

## 264 **4 Discussion**

### 265 **4.1 Variability in model response from sensitivity analysis**

266 From the different simulations performed during sensitivity analysis, there is a great amount of variability in front of the  
267 vegetation patch in wave dissipation, KE, and TKE (Figs. 4, 5 and 7). This is a result of large amount of wave dissipation and  
268 flow deceleration in front of the vegetation patch. The cross-sectional plane of the domain illustrates that the variability in KE  
269 occurs throughout the water column (Fig. 5), highlighting the 3-D impact of vegetation inputs. Interestingly, the greatest  
270 amount of variability in KE (Fig. 5) occurs at distance above the bed between 0.4 – 1.3 m at  $y = 5.6$  km, while the maximum  
271 vegetation height in all the simulations is 0.3 m. This result indicates that variation in vegetation height results in the greatest  
272 amount of KE variability above the vegetation patch. The variability in water level change is large around the lobes and behind  
273 the vegetation patch (Fig. 6). The variability in these regions is because the change in local flow velocity is adjusted by a  
274 change in water level around the vegetation patch. This is further confirmed by observing the variation in velocity profiles  
275 with depth (Fig. 8a) at a particular time instance during flood tide. Simulation 8 results in the lowest velocity of 0.06 m/s while  
276 simulation 2 results in the peak flood velocity of 0.11 m/s. Consequently, the gradient of velocity with respect to depth is  
277 greatest at all depths for simulation 8 while least for simulation 2 (Fig. 8b). The gradient reaches a maximum value at the  
278 bottom layer.

### 280 **4.2 Understanding vegetation parameterization to interpret Sobol' Indices**

281 The parameterizations involving extraction of momentum, turbulence production, and turbulence dissipation are directly  
282 affected by vegetation stem density ( $n_v$ ), diameter ( $b_v$ ), and thickness ( $t_v$ ) (Table 1). Because these mechanistic processes  
283 occur at the blade scale, the dependence on vegetation height ( $l_v$ ) is implicitly included in the parameterizations. Sobol' indices  
284 provide quantifiable information and show that vegetation height, stem density, and diameter (in decreasing order of  
285 importance) are pertinent in accurately computing kinetic energy (KE), turbulent kinetic energy (TKE), and water level. An  
286 accurate representation of KE and TKE has direct ramifications on estimating sediment transport while water level estimates  
287 can affect storm surge predictions.

288 The high sensitivity of wave dissipation to vegetation stem density highlights the need of accurate density  
289 representation to attain wave attenuation estimates, especially in open coasts. SWAN computes wave dissipation due to  
290 vegetation as a bottom layer effect. Therefore, the height of the vegetation does not affect wave dissipation to the same extent  
291 as other model outputs: KE, TKE, and water level. In addition, the equation representing the wave dissipation process in  
292 SWAN is independent of vegetation thickness, thus corresponding to the lowest Sobol' index. Vegetation thickness only  
293 appears in the turbulence dissipation term (Table 1) modifying the turbulence length scale has the least effect on any of the  
294 model responses.

### 296 **4.3 Linear curve fitting to complement Effective Quadratures based sensitivity analysis**

297 To complement the results of the Sobol' indices calculation, linear fits to the data are conducted. The main parameter  
298 contributing to the wave dissipation variability is the stem density explaining over 80% of the variability ( $R^2=0.81$ ). The least  
299 squared fit to wave dissipation that included all parameters combined explains 98% of the variability. When fitting models, it  
300 is possible to increase the explained variance by using more complex fitting models adding parameters, but doing so may result  
301 in over-fitting. The Bayesian Information Criterion (BIC; Schwarz, 1978; Aretxabaleta and Smith, 2011) provides a non-  
302 subjective metric for the best fit by penalizing over-parameterization. BIC resolves over-fitting by introducing a penalty term  
303 for the number of parameters in the model. For wave dissipation, the BIC approach identified a fit based exclusively on density  
304 as the model that best match the data while preventing over-fitting. The selection of density as the single most relevant  
305 parameter is consistent with the Sobol' indices result (Table 3).

306 The kinetic energy variability is also associated with stem density changes, but the percentage of explained variability  
307 (42%) is smaller than for wave dissipation. Diameter and height also contributed to changes in kinetic energy. The combination  
308 of stem density, height, and diameter provided the optimal fit of the data (selected by minimizing BIC) and explained 89% of  
309 variability in kinetic energy. Similar results were obtained for TKE with density explaining 45% of the variability but the  
310 combination of stem density, height, and diameter providing the optimal fit to TKE (selected by minimizing BIC) and  
311 explained 87% of the variability. The model response of water level variability is also best explained by a combination of stem  
312 density, height, and diameter (96% of water level variance explained). Thickness was not correlated with wave dissipation,  
313 kinetic energy or TKE and only contributed to water level gradient variability.

314

#### 315 **4.4 Limitations of the current sensitivity methodology**

316 The model configuration chosen includes vegetation covering a small fraction of the water column to allow for proper wave  
317 dissipation. Many species of seagrass have a larger vertical footprint and can also exhibit much higher shoot densities. The  
318 goal of the study is to provide estimates on the relative importance of the different parameters through a robust sensitivity  
319 approach. The current work assumed rigid vegetation blades, while the model is capable of including flexible vegetation by  
320 altering the blade scale. The expected effect of flexible blades would be to reduce the relative importance of vegetation length  
321 ( $l_v$ ) on the model outputs. In addition, the present work assumes a constant drag coefficient for cylindrical vegetation shape;  
322 the influence of a variability drag coefficient can be a subject of a future sensitivity study. Other model parameters such as  
323 vertical and horizontal resolution, mixing parameterization, and wave and hydrodynamic forcings will also affect model  
324 results, but were beyond the scope of the current study.

#### 325 **5 Conclusions**

326 The coupled wave-flow-vegetation module in the COAWST modeling system provides a tool to study vegetated flows in  
327 riverine, lacustrine, estuarine, and coastal environments. The resulting flow field in the presence of vegetation depends on its  
328 properties, including vegetation stem density, height, diameter, and thickness. The sensitivity of the hydrodynamic and wave  
329 conditions to changes in vegetation parameters is investigated. The sensitivity analysis helps in understanding the

330 multiparameter/multiresponse of various interactions within the model. We use an existing tool that formulates the Effective  
331 Quadratures method to quantify the sensitivity of plant input properties for the vegetation module in COAWST model. The  
332 decomposition of the variance of the model solution given by the Sobol' indices is assigned to plant parameters.

333 The method of using Sobol' indices to quantify sensitivity can be computationally expensive. One of the goals of  
334 this work is to demonstrate a robust, practical, and efficient approach for the parameter sensitivity analysis. We show that the  
335 approach of using Effective Quadratures method to select a parameter space that is consistent with physical understanding  
336 significantly reduces the computational time required to obtain the Sobol' indices.

337 The evaluation of Sobol' indices shows that the input values of plant stem density, height and to a lesser degree,  
338 diameter are consequential in determining kinetic energy, turbulent kinetic energy and water level changes. Meanwhile, the  
339 wave dissipation is mostly dependent on the variation in plant density.

340 The sensitivity analysis for the vegetation model in COAWST presented herein provides guidance for observational  
341 and modelling work by allowing future efforts to focus on constraining the most influential inputs without having to explore  
342 the entire parameter space. An accurate representation of processes causing kinetic energy and turbulent kinetic energy leads  
343 to enhanced understanding of sediment processes while accurate water level computations help to predict coastal flooding  
344 caused by storm surge. Similarly, wave attenuation measurements in open coasts are better understood with a correct  
345 representation of wave dissipation. In the future, we intend to perform a similar sensitivity analysis with the inclusion of a  
346 biological model that will affect plant growth, thus allowing a time dependence of model input and response. In addition, the  
347 influence of vegetation on sediment transport will be explored. As model complexity increases with more parameters  
348 representing additional processes, input parameter sensitivity is required for the model to be applied on practical applications.

349

## 350 **6 Code availability**

351 The Effective Quadratures methodology is an open source Python based tool designed to perform sensitivity analysis for a  
352 given physical system. The instructions to install the code along with all the open source files for this tool are detailed here:  
353 <https://github.com/Effective-Quadratures/Effective-Quadratures>. For any further inquiries about the Effective Quadratures  
354 methodology, please contact the corresponding author Dr. P. Seshadri (ps583@cam.ac.uk).

355 The COAWST model is an open source coupled hydrodynamics and wave model containing vegetation effects mainly  
356 coded in Fortran 77. This model provided the physical setting to perform the sensitivity analysis. The code is available from  
357 [https://coawstmodel-trac.sourcerepo.com/coawstmodel\\_COAWST](https://coawstmodel-trac.sourcerepo.com/coawstmodel_COAWST) after registration via email with J. C. Warner  
358 ([jcwarner@usgs.gov](mailto:jcwarner@usgs.gov)).

359

## 360 **7 Data availability**

361 The model output from various simulations used to perform sensitivity analysis in this study are available  
362 at: [http://geoport.whoi.edu/thredds/catalog/clay/usgs/users/tkalra/sensitivity\\_study/catalog.html](http://geoport.whoi.edu/thredds/catalog/clay/usgs/users/tkalra/sensitivity_study/catalog.html). The link contains a  
363 "README.txt" file that explains how the folder is organized to contain model output.

364

## 365 **8 Author contribution**

366 T. S. Kalra, and N. K. Ganju designed and simulated the numerical experiment space for sensitivity analysis. P. Seshadri  
367 developed the Effective Quadratures methodology to quantify the model sensitivity. A. Beudin provided inputs on the  
368 mechanistic processes involving the vegetation model. T. S. Kalra and A. Arextabaleta performed the data analysis from the  
369 output of sensitivity study and prepared the manuscript with contributions from all co-authors.

370

## 371 **9 Disclaimer**

372 Any use of trade, firm, or product names is for descriptive purposes only and does not imply endorsement by the U.S.  
373 Government.

374

## 375 **10 Acknowledgement**

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377 of vegetation for sensitivity studies in the early stages of work. We also thank you P. Soupy Daylander at U. S. Geological  
378 Survey, St. Petersburg Coastal and Marine Science Center for providing her feedback to improve the clarity of the paper. We  
379 thank the anonymous reviewers for their careful reading of our manuscript and their insightful comments and suggestions.

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Process	Equation	$n_v$	$l_v$	$b_v$	$t_v$
Extraction of Momentum	$F_{d,veg,u} = \frac{1}{2} C_D b_v n_v u \sqrt{u^2 + v^2}$ $C_D = \text{Plant drag coefficient}$ $u, v = \text{Horizontal velocity components at each vertical level}$	X		X	
Turbulence Production Uittenbogaard (2003)	$P_{veg} = \sqrt{(F_{d,veg,u}u)^2 + (F_{d,veg,v}v)^2}$	X		X	
Turbulence Dissipation Uittenbogaard (2003)	$D_{veg} = c_2 \frac{P_{veg}}{\tau_{eff}}$ $\tau_{eff} = \min(\tau_{free}, \tau_{veg})$ $\tau_{free} = \frac{k}{\varepsilon}$ $\tau_{veg} = \left( \frac{L^2}{c_k^2 P_{veg}} \right)^{1/3}$ $L(z) = c_l \left( \frac{1 - b_v t_v n_v}{n_v} \right)^{1/2}$ $P_{veg} = \text{Turbulence Production}$ $\tau_{free} = \text{Dissipation time scale of free turbulence}$ $k = \text{Turbulent kinetic energy}$ $\varepsilon = \text{Turbulence dissipation}$ $\tau_{veg} = \text{Dissipation time scale of free turbulence}$ $c_k = (c_\mu^0)^4 \approx 0.09$ $L = \text{typical length scale between the plants}$ $c_l = \text{Lift coefficient of order unity}$	X		X	X

<p>Wave Dissipation</p> <p>Mendez and Losada (2004)</p> <p>Dalrymple et al. (1984)</p>	$S_{d,veg} = \sqrt{\frac{2}{\pi}} g^2 \widetilde{C}_D b_v n_v \left(\frac{\tilde{k}}{\tilde{\sigma}}\right)^3$ $\frac{\sinh^3(\tilde{k}l_v) + 3 \sinh(\tilde{k}l_v)}{3\tilde{k} \cosh^3(\tilde{k}h)} \sqrt{E_{tot}} E(\sigma, \theta)$ <p><math>\widetilde{C}_D</math> = Bulk drag coefficient  <math>\tilde{k}</math> = Mean wave number  <math>\tilde{\sigma}</math> = Mean wave frequency  <math>h</math> = Water depth  <math>E_{tot}</math> = Total wave energy  <math>E</math> = Wave energy at frequency <math>\sigma</math> and direction <math>\theta</math></p>	X	X	X	
<p>Wave induced streaming</p> <p>Luhar et al., 2010</p>	$F_{s,veg} = \frac{S_{d,veg,tot} \tilde{k}}{\rho_0 \tilde{\sigma}}$ <p><math>S_{d,veg,tot}</math> = Total wave energy dissipation  <math>\tilde{k}</math> = Mean wave number  <math>\tilde{\sigma}</math> = Mean wave frequency  <math>\rho_0</math> = Reference density of seawater</p>	X	X	X	

505 **Table 1: Processes in ROMS and SWAN to model the presence of vegetation. The different input parameters (stem**  
506 **density,  $n_v$ ; height,  $l_v$ ; diameter,  $b_v$ ; and thickness,  $t_v$ ) affecting model wave and hydrodynamics are included.**  
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	Stem density (stems/m <sup>2</sup> ) $n_v$	Height (m) $l_v$	Diameter (mm) $b_v$	Thickness (mm) $t_v$
1.	144.3	0.24	6.0	0.6
2.	62.1	0.17	3.0	0.6
3.	226.5	0.3	6.0	0.3
4.	144.3	0.17	9.0	0.9
5.	144.3	0.3	3.0	0.9
6.	62.1	0.3	6.0	0.3
7.	226.5	0.17	6.0	0.3
8.	144.3	0.3	9.0	0.9
9.	144.3	0.24	9.0	0.3
10.	226.5	0.24	6.0	0.9
11.	144.3	0.24	3.0	0.3
12.	62.1	0.24	6.0	0.9
13.	62.1	0.17	6.0	0.3
14.	62.1	0.24	9.0	0.6
15.	144.3	0.17	3.0	0.9

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**Table 2: Plant property input combination for different simulations during sensitivity analysis.**

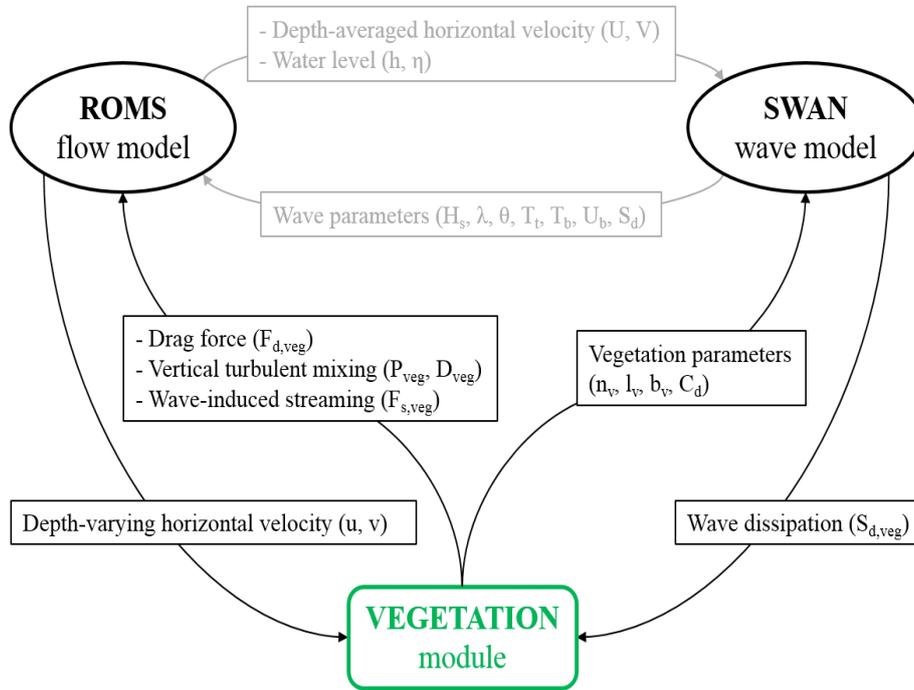
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	Plant Stem Density $n_v$	Plant Height $l_v$	Plant Diameter $b_v$	Plant Thickness $t_v$
Wave dissipation	0.68	0.24	0.032	0.01
Kinetic energy	0.36	0.44	0.12	0.03
Maximum water level change	0.38	0.43	0.15	0.01
Turbulent kinetic energy	0.35	0.42	0.12	0.03

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**Table 3: Sobol Indices for all the outputs.**

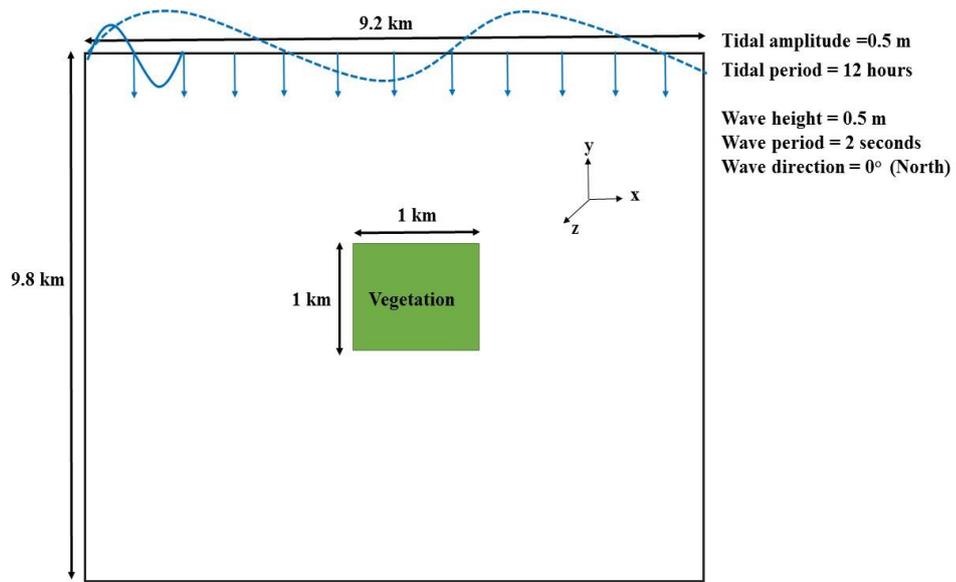
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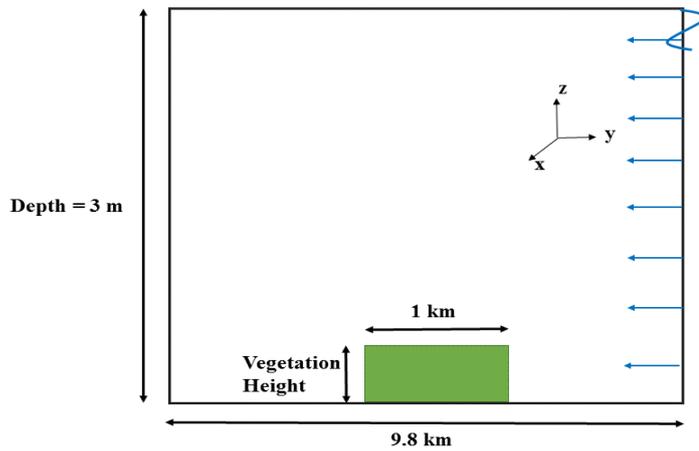
576 **Figure 1: Schematic showing the vegetation module implementation in COAWST model (Figure adapted from Beudin**  
577 **et. al, 2017)**

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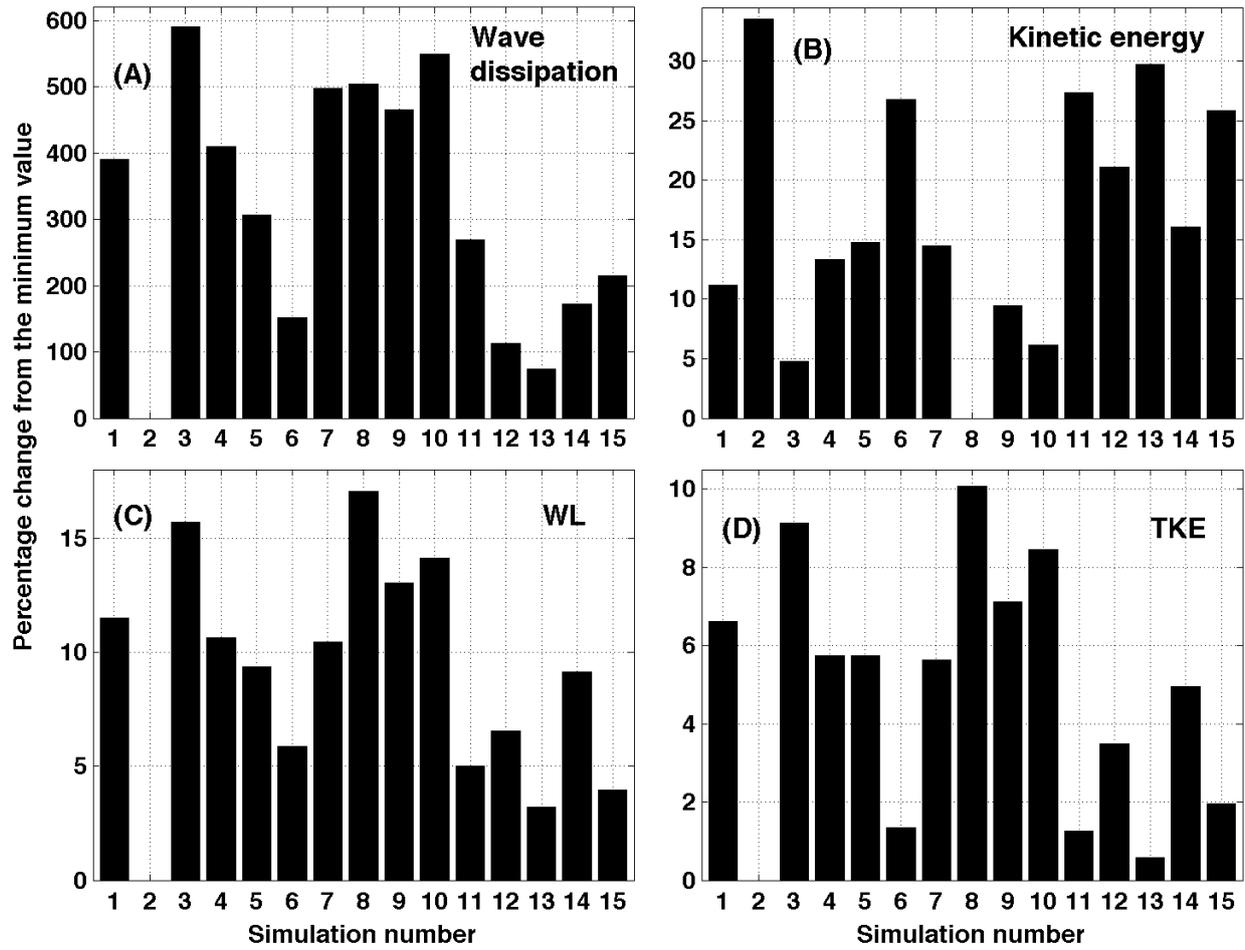
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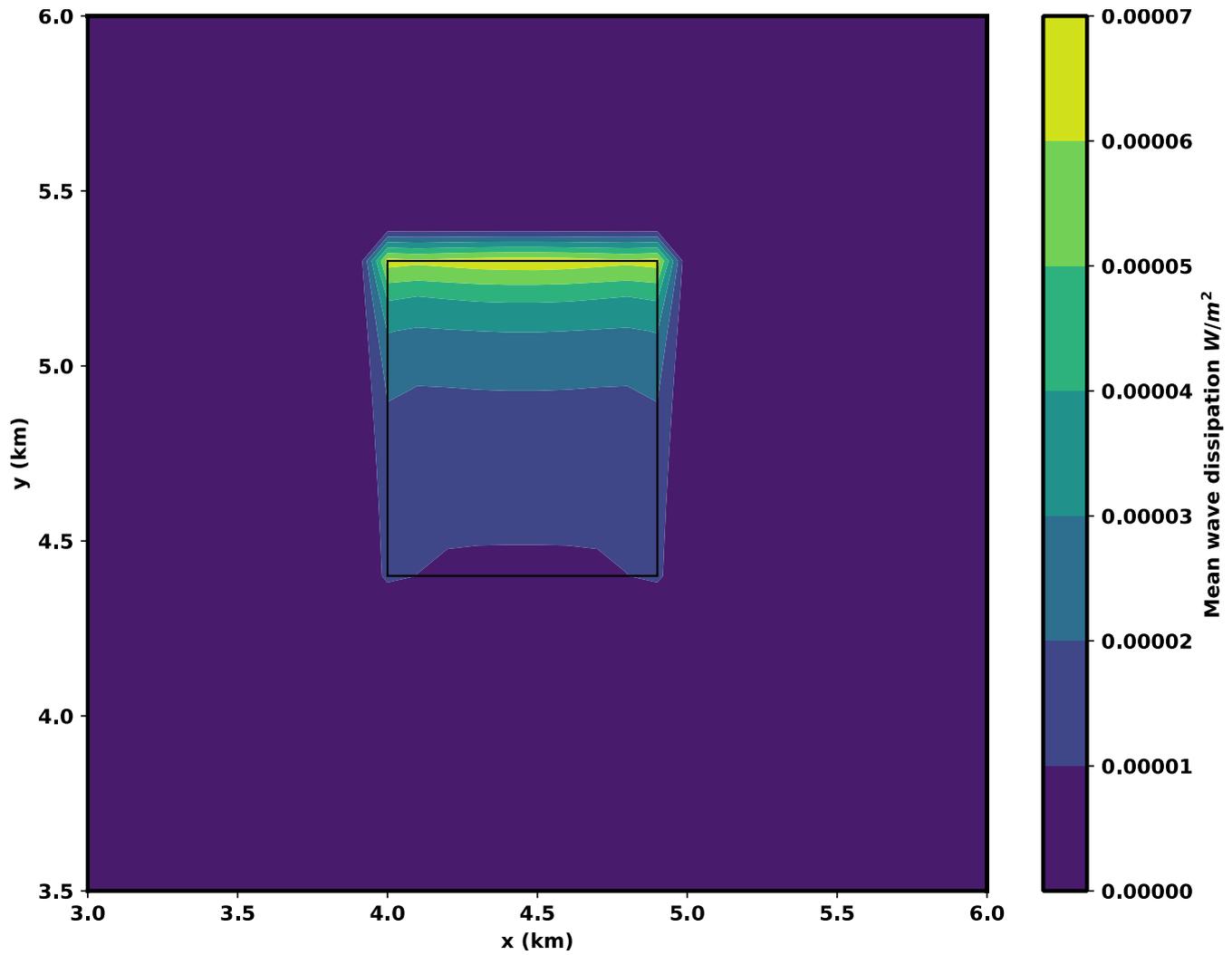
Figure 2: Schematic showing the idealized domain (Not drawn to scale)

(a) Plan view and (b) Cross-sectional view.



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 588 **Figure 3: Percentage change from minimum for the four impact parameters: a) wave dissipation; b) kinetic energy;**  
 589 **c) maximum water level change (WL); and d) Turbulent Kinetic Energy (TKE).**

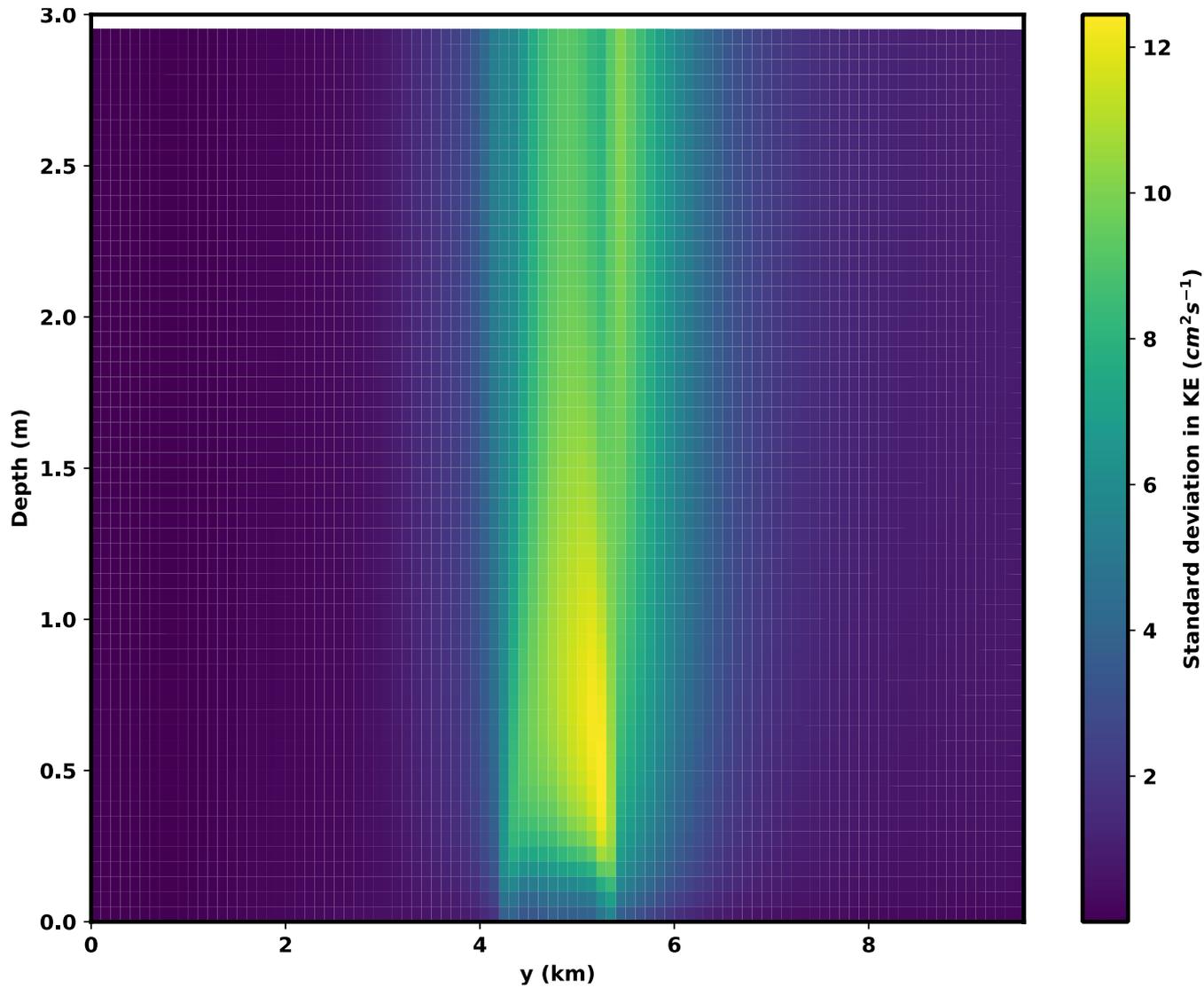
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609 **Figure 4: Standard deviation from wave dissipation ( $W\ m^{-2}$ ) in presence of vegetation (Plan view). The area of the**  
610 **vegetation patch is highlighted in the middle of the domain.**

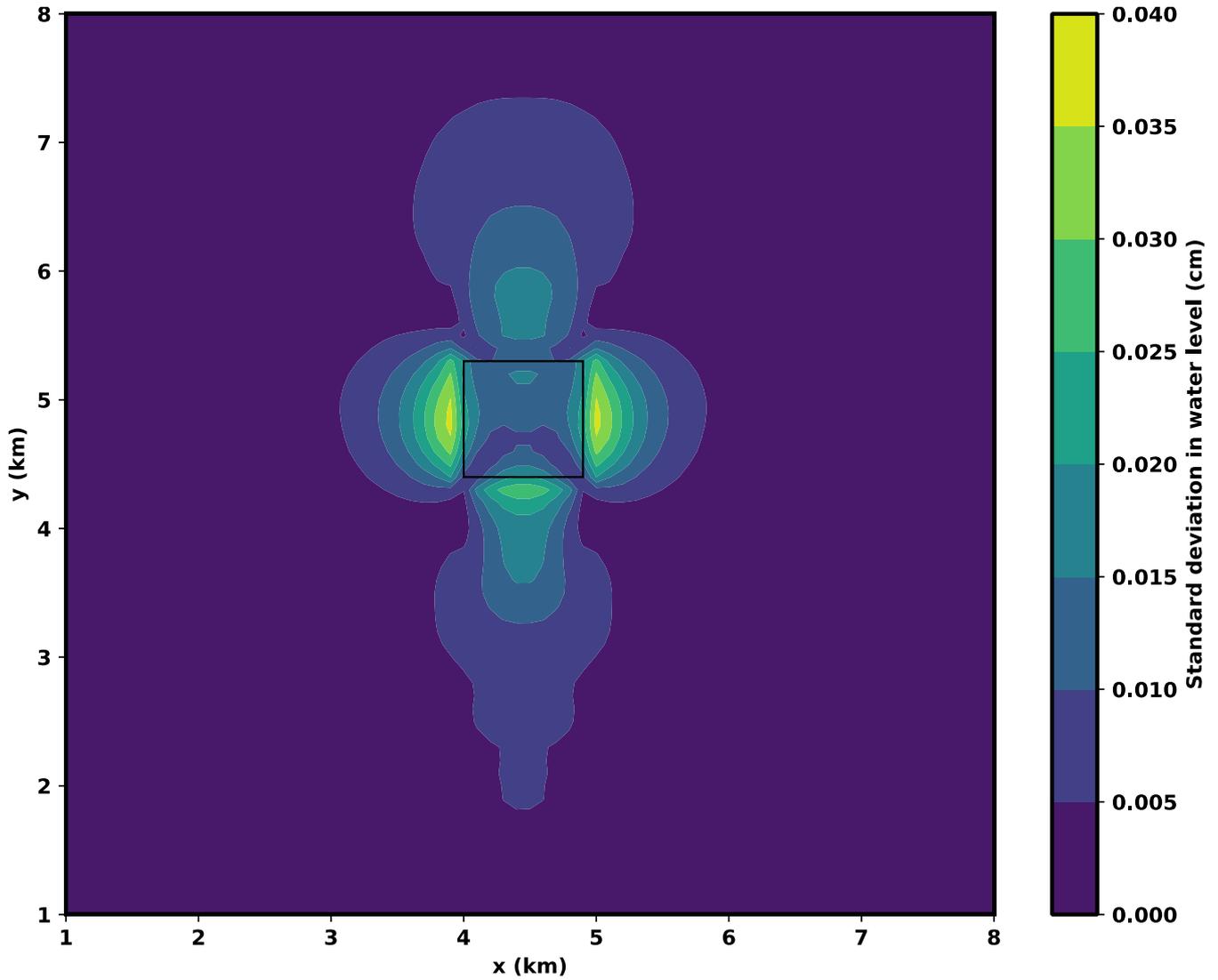
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Figure 5: Standard deviation in kinetic energy ( $cm^2 s^{-1}$ ) in presence of vegetation (cross-sectional view).

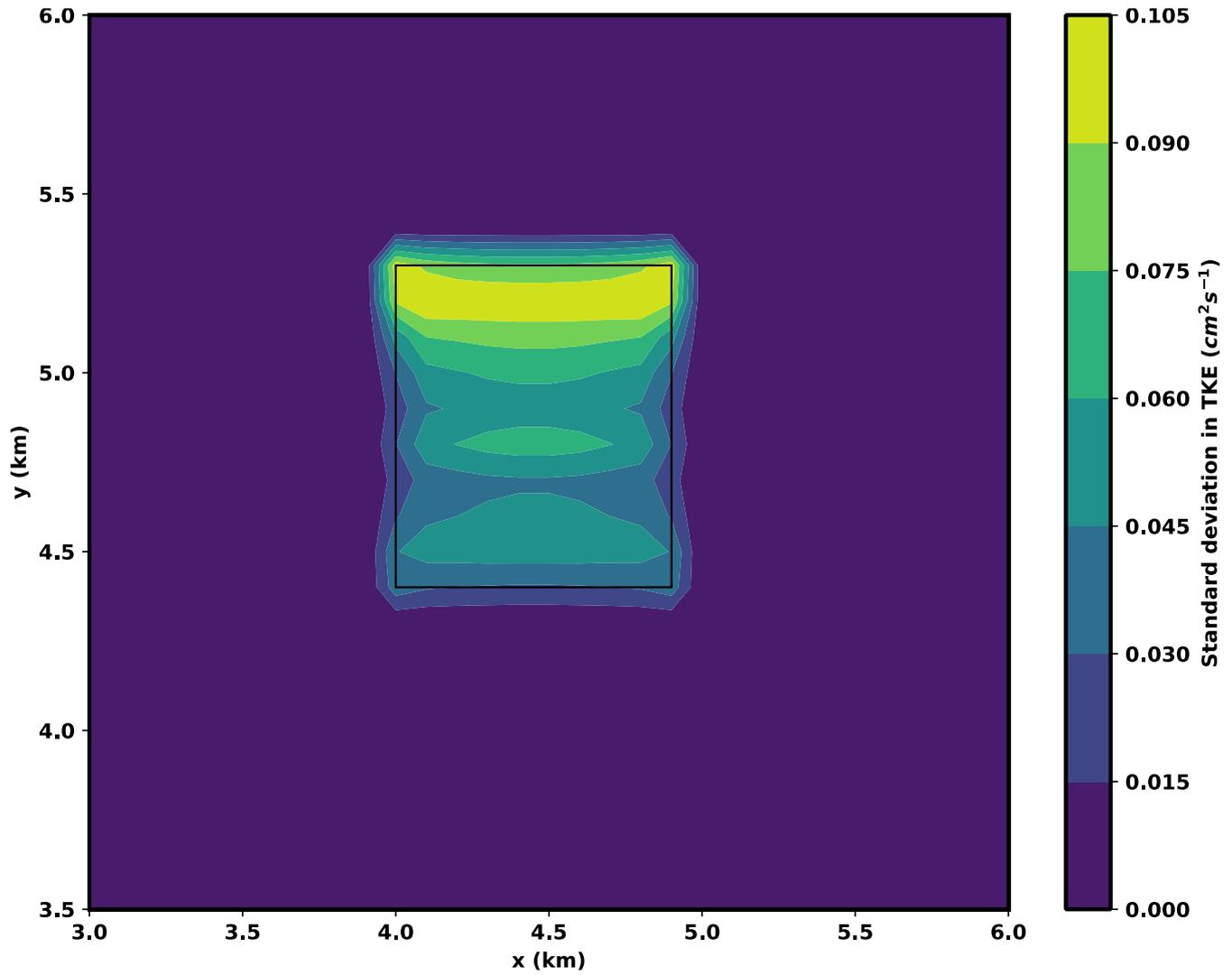
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**Figure 6: Standard deviation in water level in presence of vegetation (plan view). The area of the vegetation patch is highlighted in the middle of the domain.**

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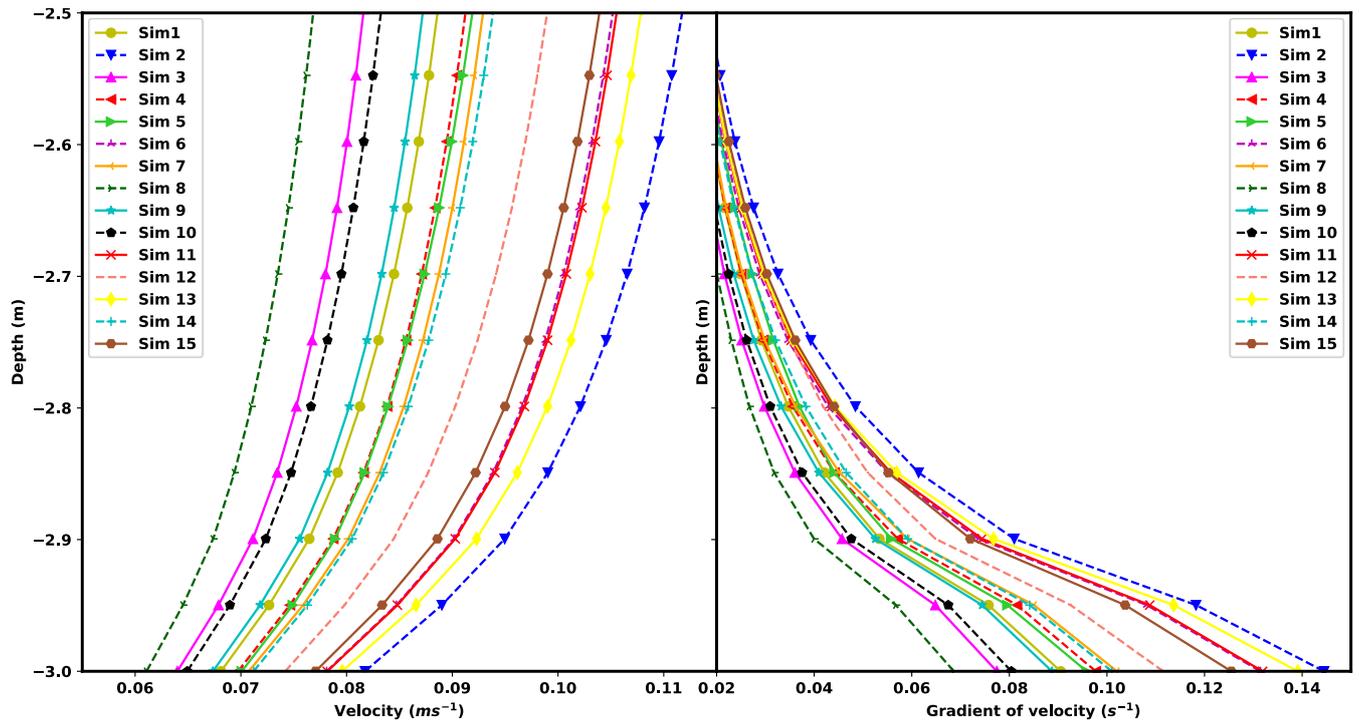
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**Figure 7: Standard deviation in Turbulent Kinetic Energy (TKE) ( $\text{cm}^2 \text{s}^{-1}$ ) in presence of vegetation (plan view). The area of the vegetation patch is highlighted in the middle of the domain.**



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661 **Figure 8: (a) Velocity ( $\text{m s}^{-1}$ ) profile and (b) Vertical gradient ( $\text{s}^{-1}$ ) of velocity profile varying with depth in front of the**  
662 **vegetation patch at a particular time instance during flood for different simulations.**

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