

The Cloud Feedback Model Intercomparison Project (CFMIP) contribution to CMIP6.

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Abstract

The primary objective of CFMIP is to inform future assessments of cloud feedbacks through improved understanding of cloud-climate feedback mechanisms and better evaluation of cloud processes and cloud feedbacks in climate models. However, the CFMIP approach is also increasingly being used to understand other aspects of climate change, and so a second objective has now been introduced, to improve understanding of circulation, regional-scale precipitation, and non-linear changes. CFMIP is supporting ongoing model inter-comparison activities by coordinating a hierarchy of targeted experiments for CMIP6, along with a set of cloud related output diagnostics. CFMIP contributes primarily to addressing the CMIP6 questions "How does the Earth System respond to forcing?" and "What are the origins and consequences of systematic model biases?" and supports the activities of the WCRP Grand Challenge on Clouds, Circulation and Climate Sensitivity.

A compact set of Tier 1 experiments is proposed for CMIP6 to address the question: "1) What are the physical mechanisms underlying the range of cloud feedbacks and cloud adjustments predicted by climate models, and which models have the most credible cloud feedbacks?" Additional Tier 2 experiments are proposed to address the following questions: 2) Are cloud feedbacks consistent for climate cooling and warming, and if not, why? 3) How do cloud-radiative effects impact the structure, the strength and the variability of the general atmospheric circulation in present and future climates? 4) How do responses in the climate system due to changes in solar forcing differ from changes due to CO₂, and is the response sensitive to the sign of the forcing? 5) To what extent is regional climate change per CO₂ doubling state-dependent (nonlinear), and why? 6) Are climate feedbacks during the 20th century different to those acting on long term climate change and climate sensitivity? 7) How do regional climate responses (e.g. in precipitation) and their uncertainties in coupled models arise from the combination of different aspects of CO₂ forcing and sea surface warming?

CFMIP also proposes a number of additional model outputs in the CMIP DECK, CMIP6 Historical and CMIP6 CFMIP experiments, including COSP simulator outputs and process diagnostics to address the following questions: 1) How well do clouds and other relevant variables simulated by models agree with observations? 2) What physical processes and mechanisms are important for a credible simulation of clouds, cloud feedbacks and cloud adjustments in climate models? 3) Which models have the most credible representations of processes relevant to the simulation of clouds? 4) How do clouds and their changes interact with other elements of the climate system?

1 Introduction

Inter-model differences in cloud feedbacks continue to be the largest source of uncertainty in predictions of equilibrium climate sensitivity (Boucher et al., 2013). Although the ranges of cloud feedbacks and climate sensitivity from comprehensive climate models have not reduced in recent years, considerable progress has been made in understanding (a) which types of clouds contribute most to this spread (e.g. Bony and Dufresne 2005; Webb et al., 2006; Zelinka et al., 2013), (b) the role of cloud adjustments in climate sensitivity (e.g. Gregory and Webb, 2008; Andrews and Forster, 2008; Kamae and Watanabe, 2012; Zelinka et al., 2013), (c) the processes and mechanisms which are (and are not) implicated in cloud

56 feedbacks, both in fine resolution models (e.g. Rieck et al., 2012; Bretherton et al., 2015) and in comprehensive climate
57 models (e.g. Brient and Bony 2012; Sherwood et al., 2014; Zhao, 2015; Webb et al., 2015b), (d) the inconstancy of cloud
58 feedbacks and effective climate sensitivity (e.g. Senior and Mitchell, 2000; Williams et al., 2008; Andrews et al., 2012;
59 Geoffroy et al., 2013; Armour et al., 2013; Andrews and Gregory, 2016) and (e) the extent to which models with stronger or
60 weaker cloud feedbacks or climate sensitivities agree with observations (e.g. Fasullo and Trenberth, 2012; Su et al., 2014; Qu
61 et al., 2014; Sherwood et al., 2014; Myers and Norris, 2016). Additionally, our ability to evaluate model clouds using
62 satellite data has benefited from the increasing use of satellite simulators. This approach, first introduced by Yu et al, 1996
63 for use with data from the International Satellite Cloud Climatology Project (ISCCP) attempts to reproduce what a satellite
64 would observe given the model state. Such approaches enable more quantitative comparisons to the satellite record (e.g. Yu
65 et al., 1996; Klein and Jakob, 1999; Webb et al.; 2001; Bodas-Salcedo et al., 2011; Cesana and Chepfer, 2013). Much of our
66 improved understanding in these areas would have been impossible without the continuing investment of the scientific
67 community in successive phases of the Coupled Model Intercomparison Project (CMIP), and its co-evolution in more recent
68 years with the Cloud Feedback Model Intercomparison Project (CFMIP).

69 CFMIP started in 2003 and its first phase (CFMIP-1) organised an intercomparison based on perpetual July SST forced
70 Cess style +2K experiments and 2xCO₂ equilibrium mixed-layer model experiments containing ISCCP simulator in parallel
71 with CMIP3 (McAvaney and Le Treut, 2003). CFMIP-1 had a substantial impact on the evaluation of clouds in models and
72 in the identification of low level cloud feedbacks as the primary cause of inter-model spread in cloud feedback, which
73 featured prominently in the fourth and fifth IPCC assessments (Randall et al., 2007; Boucher et al., 2013).

74 The subsequent objective of CFMIP-2 was to inform improved assessments of climate change cloud feedbacks by
75 providing better tools to support evaluation of clouds simulated by climate models and understanding of cloud-climate
76 feedback processes. CFMIP-2 organized further experiments as part of CMIP5 (Bony et al., 2011; Taylor et al., 2012),
77 introducing seasonally varying SST perturbation experiments for the first time, as well as fixed SST CO₂ forcing experiments
78 to examine cloud adjustments. CFMIP-2 also introduced idealized ‘aquaplanet’ experiments into the CMIP family of
79 experiments. These experiments were motivated by extensive research in the framework of the aqua-planet experiment
80 (Neale and Hoskins, 2000, Blackburn and Hoskins, 2013) and the particular finding, based on a small subset of models, that
81 the global mean cloud feedback of more realistic model configurations could be reproduced, and more easily investigated,
82 using the much simpler aqua-planet configuration (Medeiros et al., 2008). CFMIP-2 proposed the inclusion of the abrupt
83 CO₂ quadrupling AOGCM (atmosphere–ocean general circulation model) experiment in the core experiment set of CMIP5,
84 based on the approach of Gregory et al., 2004, which subsequently formed the basis for equilibrium climate sensitivity
85 estimates from AOGCMs (Andrews et al., 2012). Additionally CFMIP-2 introduced satellite simulators to CMIP via the
86 CFMIP Observation Simulator Package (COSP, Bodas-Salcedo et al., 2011); not only the ISCCP simulator, but additional
87 simulators to facilitate the quantitative evaluation clouds using a new generation of active radars and lidars in space. CFMIP-
88 2 also introduced into CMIP5 process diagnostics such as temperature and humidity budget tendency terms and high
89 frequency ‘cfSites’ outputs at 120 locations around the globe. In an effort less directly connected to CMIP, CFMIP organized
90 a joint project with the GEWEX Global Atmospheric System Study (GASS) called CGILS (the CFMIP-GASS
91 Intercomparison of LES and SCMs) to develop cloud feedback intercomparison cases to assess the physical credibility of
92 cloud feedbacks in climate models by comparing Single Column Model (SCM) versions of General Circulation Models
93 (GCMs) with high resolution Large Eddy Simulations (LES) models. CFMIP-2 also developed the CFMIP-OBS data portal
94 and the CFMIP diagnostic codes catalogue. For more details, and for a full list of CFMIP related publications, please refer
95 to the CFMIP website (<http://www.earthsystemcog.org/projects/cfmip>).

96 Studies arising from CFMIP-2 include numerous single and multi-model evaluation studies which use COSP to make
97 quantitative and fair comparisons with a range of satellite products (e.g. Kay et al., 2012; Franklin et al., 2013; Klein et al.,
98 2013, Lin et al., 2014, Chepfer et al., 2014.). COSP has also enabled studies attributing cloud feedbacks and cloud
99 adjustments to different cloud types (e.g. Zelinka et al., 2013; Zelinka et al., 2014; Tsushima et al., 2015). CFMIP-2
100 additionally enabled the finding that idealized ‘aquaplanet’ experiments without land, seasonal cycles or Walker circulations
101 are able to reproduce the essential differences between models’ global cloud feedbacks and cloud adjustments in a substantial
102 ensemble of models (Ringer et al., 2014; Medeiros et al., 2015). Process outputs from CFMIP have also been used to develop
103 and test physical mechanisms proposed to explain and constrain inter-model spread in cloud feedbacks in the CMIP5 models
104 (e.g. Sherwood et al., 2014; Brient et al., 2015; Webb et al., 2015a; Nuijens et al., 2015a,b; Dal Gesso at al., 2015). CGILS
105 has demonstrated a consensus in the responses of LES models to climate forcings and identified shortcomings in the physical
106 representations of cloud feedbacks in climate models (e.g. Blossey et al., 2013; Zhang et al., 2013; Dal Gesso at al., 2015).
107 The CFMIP experiments have additionally formed the basis for coordinated experiments to explore the impact of cloud
108 radiative effects on the circulation (Stevens et al., 2012; Fermepin and Bony 2014; Crueger and Stevens 2015; Li et al., 2015;
109 Harrop and Hartmann 2016), the impact of parametrized convection on cloud feedback (Webb et al., 2015b) and the
110 mechanisms of negative shortwave cloud feedback in mid to high latitudes (Ceppi et al., 2015). Additionally the CFMIP
111 experiments have, due to their idealized nature, proven useful in a number of studies not directly related to clouds, but instead
112 analyzing the responses of regional precipitation and circulation patterns to CO₂ forcing and climate change (e.g. Bony et al.,
113 2013; Chadwick et al., 2014; He and Soden 2015; Oueslati et al., 2016). Studies using CFMIP-2 outputs from CMIP5 remain
114 ongoing and further results are expected to feed into future assessments of the representation of clouds and cloud feedbacks in
115 climate models.

116 The primary objective of CFMIP is to inform future assessments of cloud feedbacks through improved understanding of
117 cloud-climate feedback mechanisms and better evaluation of cloud processes and cloud feedbacks in climate models.
118 However, the CFMIP approach is also increasingly being used to understand other aspects of climate change, and so a second
119 objective has been introduced, to improve understanding of circulation, regional-scale precipitation, and non-linear changes.

120 This involves bringing climate modelling, observational and process modelling communities closer together and providing
121 better tools and community support for evaluation of clouds and cloud feedbacks simulated by climate models and for
122 understanding of the mechanisms underlying them. This is achieved by:

- 123 • Coordinating model inter-comparison activities which include experimental design as well as specification of
124 model output diagnostics to support quantitative evaluation of modelled clouds with observations (e.g. COSP)
125 and in-situ measurements (e.g. cfSites) as well as process-based investigation of cloud maintenance and
126 feedback mechanisms (e.g. cfSites, temperature and humidity tendency terms)
- 127 • Developing and improving support infrastructure including COSP, CFMIP-OBS and the CFMIP diagnostic
128 codes catalogue.
- 129 • Fostering collaboration with the observational and cloud process modelling communities via annual CFMIP
130 meetings and international funded projects.

132 This paper describes and documents the CFMIP contribution to the current phase on the Coupled Model Intercomparison
133 Project (CMIP6, Eyring et al., 2016). It is anticipated that CFMIP-3 will eventually be broader than what is described here,
134 for instance including studies with process models, but for the purposes of this document CFMIP-3 should be considered to
135 be synonymous with the CFMIP contribution to CMIP6. CFMIP-3 touches, to differing degrees, on each of the three
136 questions around which CMIP6 is organized. With its focus on cloud feedback, CFMIP-3 is central to CMIP6's attempt to
137 answer the question: 'How does the Earth system respond to forcing?' But as illustrated in the remainder of this document,
138 CFMIP-3 also offers the opportunity to contribute to the other two guiding questions of CMIP6. Through its strong model
139 evaluation component it stands to help answer the question: 'What are the origins and consequences of systematic model
140 biases?' CFMIP-3 will also help answer the question: 'How can we assess future climate changes given climate variability,
141 climate predictability, and uncertainties in scenarios?' For example the *amip-piForcing* experiment proposed below will
142 support studies relating cloud variability and feedbacks on observable timescales to long term cloud feedbacks (Andrews,
143 2014; Gregory and Andrews, 2016).

144 The CFMIP-3 experiments proposed for CMIP6 are outlined below in Section 2. Section 3 describes the diagnostics
145 outputs proposed by CFMIP for the CFMIP-3 experiments and other experiments within CMIP. We provide a summary of
146 the CFMIP-3 contribution to CMIP6 in Section 5.

148 2 CFMIP-3 Experiments

149 The CFMIP-3 experiments are summarised in Figure 1 and Tables 1 and 2, and are described in detail below. Most of the
150 CFMIP-3 experiments are based on CO₂ concentration forced *amip*, *piControl* and *abrupt-4xCO2* CMIP DECK (Diagnostic,
151 Evaluation and Characterization of Klima) experiments (Eyring et al., 2016). Unless otherwise specified below, the
152 CFMIP-3 experiments should be configured consistently with the DECK experiments on which they are based, using
153 consistent model formulation, and forcings and boundary conditions as specified by Eyring et al., 2016. Following the
154 CMIP6 design protocol, groups of experiments are motivated by science questions and are separated into Tiers 1 and 2
155 (Eyring et al., 2016). It is a requirement for participation by modelling groups in the CFMIP-3/CMIP6 model
156 intercomparison that all Tier 1 experiments be performed and published through the ESGF, so as to support CFMIP's Tier 1
157 science question. Tier 2 experiments are optional, and are associated with additional science questions. Any subset of Tier
158 2 experiments may be performed. All model output archived by CFMIP/CMIP6 is expected to be made available under the
159 same terms as CMIP output. Most modelling groups currently release their CMIP data for unrestricted use. Our analysis
160 plans for the CFMIP-3 experiments are summarised in Appendix A.

162 2.1 CFMIP-3 Tier 1 Experiments

163 Lead coordinator: Mark Webb

164 Science Question: What are the physical mechanisms underlying the range of cloud feedbacks and cloud adjustments
165 predicted by climate models, and which of the cloud responses are the most credible?

166 Equilibrium climate sensitivity (ECS) can be estimated using an idealized AOGCM experiment such as the *abrupt-4xCO2*
167 experiment in the CMIP6 DECK, at the same time statistically separating the global mean contributions from climate
168 feedbacks and adjusted radiative forcing due to CO₂ (Gregory et al. 2004, Andrews et al., 2012). However understanding the
169 physical processes underlying cloud feedbacks and adjustments requires diagnosis in SST forced experiments with
170 atmosphere-only general circulation models (AGCMs), which can resolve cloud feedbacks and adjustments independently
171 from each other and with minimal statistical noise at regional scales, while faithfully reproducing the inter-model differences
172 in global values from the fully coupled models (Ringer et al., 2014). (The ability of these AGCM experiments to reproduce
173 the inter-model differences in global cloud feedbacks and adjustments from coupled models indicates that they do not
174 strongly depend on different ocean model formulations or SST biases). The CFMIP-2 *amip4xCO2* experiments in CMIP5,
175 which quadrupled CO₂ while leaving SSTs at present day values (Bony et al., 2011), allowed the land/tropospheric
176 adjustment process and the cloud adjustment to CO₂ to be examined in this way for the first time in the multi-model context
177 (Kamae and Watanabe, 2012; Ringer et al., 2014; Kamae et al. 2015) in conjunction with the CMIP5 *sstClim/sstClim4xCO2*
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179
180

181 experiments which were based on climatological preindustrial SSTs (Andrews et al., 2012; Zelinka et al., 2013; Vial et al.,
182 2013). These experiments have additionally formed the basis for more in-depth studies with individual models (e.g. Wyant et
183 al., 2012; Kamae and Watanabe, 2013; Bretherton et al., 2014, Ogura et al., 2014). The CFMIP-2/CMIP5 *amip4K* and
184 *amipFuture* SST perturbed atmosphere-only experiments (Bony et al., 2011) have been used to examine cloud feedbacks in
185 greater detail (e.g. Brient and Bony, 2012; Bretherton et al., 2014; Lacagnina et al., 2014; Bellomo and Clement, 2015; Webb
186 et al., 2015b), often in conjunction with simulator outputs (e.g. Gordon and Klein, 2014; Chepfer et al., 2014; Tsushima et al.,
187 2015, Ceppi et al., 2016) and CFMIP process diagnostics (e.g. Webb and Lock, 2013; Sherwood et al., 2014; Brient et al.,
188 2015; Webb et al., 2015a; Dal Gesso et al., 2015). Similarly, these experiments have been used to investigate regional
189 responses of various quantities to direct radiative forcing due to increasing CO₂ concentrations and/or increases in SST,
190 including precipitation (e.g. Ma and Xie, 2013; Huang et al., 2013; Widlansky et al., 2013; Kent et al., 2015; Long et al.,
191 2016), circulation (e.g. He et al., 2014; Zhou et al., 2014; Kamae et al., 2014; Bellomo and Clement, 2015; Shaw and Voigt,
192 2015) and stability (e.g. Qu et al., 2015).

193 A more idealized set of fixed SST experiments proposed by CFMIP-2 for CMIP5 (*aquaControl*, *aqua4xCO2*, and
194 *aqua4K*) based on zonally symmetric, fixed season ‘aquaplanet’ configurations without land have been shown to reproduce
195 the inter-model differences in global mean cloud adjustments and feedbacks from realistic experiments surprisingly
196 effectively (Medeiros et al., 2008; Ringer et al., 2014; Medeiros et al., 2015) as well as many aspects of the zonal mean
197 circulation response (Medeiros et al., 2015). This indicates that those features of the climate system excluded from these
198 experiments (i.e. the ocean, land, seasonal cycle, monsoon and Walker circulations) are not central to understanding inter-
199 model differences in global mean cloud feedbacks and adjustments, and demonstrates the value of aquaplanet experiments for
200 investigating the origin of such differences, as well as differences in zonally averaged precipitation and circulation and their
201 responses to climate change (e.g. Stevens et al., 2012; Bony et al., 2013; Oueslati and Bellon, 2013; Fermepin and Bony
202 2014; Voigt and Shaw 2015). The aquaplanet experiments have the benefit not only of being less computationally expensive
203 than alternative experiments (requiring only 5-10 years to get a robust signal); they are also much more straightforward to
204 analyse, as their behaviour can mostly be characterized by examining zonal means, avoiding the analysis overhead of
205 compositing which is generally required in realistic model configurations to isolate the various cloud regimes. Aqua-planet
206 simulations (and other idealized) experiments are particularly effective at highlighting model differences, for instance in the
207 placement of the tropical rain bands, or in the representation of cloud changes with warming, as it is not possible to tune them
208 to observations in the same way as is for more realistic configurations (e.g., Stevens and Bony, 2013).

209 The CMIP5/CFMIP-2 experiments and diagnostic outputs have thus enabled considerable progress on a number of
210 questions. However, participation by a larger fraction of modelling groups is desired in CMIP6 to enable a more
211 comprehensive assessment of the uncertainties across the full multi-model ensemble. Our proposal is therefore to retain the
212 CFMIP-2/CMIP5 experiments (known in CMIP5 as *amip4K*, *amip4xCO2*, *amipFuture*, *aquaControl*, *aqua4xCO2* and
213 *aqua4K*) in Tier 1 for CFMIP/CMIP6. These are summarised in Table 1 (the names have been changed slightly compared to
214 the CMIP5 equivalents to fit in with a wider naming convention of CMIP6). The set up for each of these experiments is
215 described below. (For output requirements from these and other experiments please refer to Section 3).

216 *amip*: This is a single ensemble member of the CMIP DECK *amip* experiment which contains additional outputs which are
217 required both for model evaluation using COSP, and for interpretation of feedbacks and adjustments in conjunction with the
218 *amip-p4K*, *amip-4xCO2*, *amip-future4K* and *amip-m4K* experiments.

219 *amip-p4K* (formerly *amip4K*): The same as the *amip* DECK experiment, except that SSTs are subject to a uniform
220 warming of 4K. This warming should be applied to the ice free ocean surface only. Sea ice and SSTs in grid boxes
221 containing sea ice remain the same as in the *amip* DECK experiment.

222 *amip-future4K* (formerly *amipFuture*): The same as the *amip* DECK experiment, except that a composite SST warming
223 pattern derived from the CMIP3 coupled models is added to the AMIP SSTs (see Appendix C for details). As with the *amip-*
224 *p4K* experiment, the warming pattern should only be applied to the ice free ocean surface, and sea ice and SSTs in grid boxes
225 containing sea ice should remain the same as in the *amip* DECK experiment. The warming pattern should be scaled to ensure
226 that the global mean SST increase averaged over the ice free oceans is 4K. Care should be taken to ensure that SSTs are
227 increased in any inland bodies of water and near coastal edges, for example by linearly interpolating the provided warming
228 pattern dataset to fill in missing data before re-gridding to the target resolution.

229 *amip-4xCO2* (formerly *amip4xCO2*): The same as the *amip* experiment within the DECK, except that the CO₂
230 concentration seen by the radiation scheme is quadrupled. The CO₂ seen by the vegetation should be the same as in the *amip*
231 DECK experiment. This experiment gives an indication of the adjusted radiative forcing due to CO₂ quadrupling, including
232 stratospheric, land surface, tropospheric and cloud adjustments.

233 The configuration of the *aqua-control*, *aqua-p4K* and *aqua-4xCO2* experiments are unchanged compared to their
234 equivalents in CFMIP-2/CMIP5, except that the simulation length has been extended to 10 years to improve the signal to
235 noise ratio. Further details of their experimental set up are included in Appendix B.

236 We also propose to use the Tier 1 experiments as the foundation for further experiments planned in the context of the
237 Grand Challenge on Clouds, Circulation and Climate Sensitivity (Bony et al., 2015). These will include for example
238 sensitivity experiments to assess the impacts of different physical processes on cloud feedbacks and regional
239 circulation/precipitation responses and also to test specifically proposed cloud feedback mechanisms (e.g. Webb et al., 2015b,
240 Ceppi et al., 2015). Additional experiments further idealizing the aquaplanet framework to a non-rotating rotationally
241 symmetric case are also under development (e.g. Popke et al., 2013). These will be proposed as additional Tier 2
242 experiments at a future time, or coordinated by CFMIP outside of CMIP6.

243 2.2 amip minus 4K Experiment (Tier 2)

244 Lead Coordinators: Mark Webb and Bjorn Stevens

245

246 Science Question: Are cloud feedbacks consistent for climate cooling and warming, and if not, why?

247

248 There is some evidence to suggest that cloud feedbacks might operate differently in response to cooling rather than warming.
249 For example, Yoshimori et al., 2009 found a positive shortwave cloud feedback in a CO₂ doubling experiment with a
250 particular GCM, but noted a tendency for it to become weaker or even negative in cooling experiments designed to replicate
251 the climate of the last glacial maximum. They suggested that this might be related to different displacements of mixed-phase
252 clouds in the two scenarios. For small enough changes where linearity is a good approximation, one would expect the cloud
253 response to cooling and warming to be the same, differing only in sign, resulting in an identical cloud feedback expressed per
254 degree of global temperature change. But for larger perturbations this symmetry of response may no longer hold. A
255 warming or cooling of the atmosphere of equal magnitude while maintaining relative humidity will for example generate
256 different changes in absolute humidity, and its horizontal and vertical gradients, which have been linked to cloud feedbacks
257 (Brient and Bony, 2013; Sherwood et al., 2014), the atmospheric lapse rate and circulation which influences clouds and
258 depends in part on the absolute humidity (Held and Soden, 2006; Qu et al., 2015) and additionally on extratropical cloud
259 optical depth feedbacks which may be related to adiabatic cloud liquid water contents (Gordon and Klein, 2014) or phase
260 changes that depend upon whether a given volume crosses the 0 degree isotherm in the climate change (Ceppi et al. 2015).

261 The configuration of the *amip-m4K* experiment will be the same as the *amip-p4K* experiment, except that the sea surface
262 temperatures are uniformly reduced by 4K rather than increased. This cooling should be applied to sea ice free grid boxes
263 only. Sea ice and SSTs in grid boxes containing sea ice should remain the same as in the *amip* DECK experiment. In models
264 which employ a fixed lower threshold near freezing for the SST used in the calculation of the surface fluxes, this should
265 ideally also be reduced by 4K. This experiment will contain CFMIP COSP and process outputs so as to support the
266 investigation of inconsistent responses of clouds to a cooling vs. a warming climate in a controlled way through comparison
267 with the *amip-p4K* experiment. This experiment also complements the abrupt 0.5xCO₂ and the -4% solar experiments in that
268 one can identify asymmetries in the warming/cooling response with and without interactions with the ocean. As such we hope
269 that these experiments will provide useful synergies with the Palaeoclimate Model Intercomparison Project (PMIP) CMIP6
270 experiments (Kageyama et. al., 2016), for example in interpreting differing cloud feedbacks between future CO₂ forced
271 experiments and those representing the Last Glacial Maximum, as highlighted by Yoshimori et al., 2009.
272

273 2.3 Atmosphere-only experiments without longwave cloud radiative effects. (Tier 2)

274 Lead Coordinators: Sandrine Bony and Bjorn Stevens

275

276 Science question: How do cloud-radiative effects impact the structure, the strength and the variability of the general
277 atmospheric circulation in present and future climates?

278

279 It is increasingly recognized that clouds, and atmospheric cloud-radiative effects in particular, play a critical role in the
280 general circulation of the atmosphere and its response to global warming or other perturbations: they have been found to
281 modulate the structure, the position and shifts of the ITCZ (e.g. Slingo and Slingo 1988; Randall et al., 1989; Sherwood et al
282 1994; Bergman and Hendon 2000; Hwang and Frierson, 2013; Fermepin and Bony 2014; Voigt et al., 2014; Loeb et al.,
283 2015), the organisation of convection in tropical waves, Madden-Julian Oscillations and other forms of convective
284 aggregation (e.g. Lee et al., 2001; Lin and Mapes, 2004; Bony and Emanuel, 2005; Zurovac-Jevtic et al., 2006; Crueger and
285 Stevens, 2015; Muller and Bony, 2015), the extra-tropical circulation and the position of eddy-driven jets (e.g. Ceppi et al.,
286 2012; Ceppi et al., 2014; Grise and Polvani 2014; Li et al., 2015; Voigt and Shaw, 2015), and modes of interannual to
287 decadal climate variability (e.g. Bellomo et al., 2015; Rädel et al., 2016; Yuan et al., 2016). A better assessment of this role
288 would greatly help to interpret model biases (how much do biases in cloud-radiative properties contribute to biases in the
289 structure of the ITCZ, in the position and strength of the storm tracks, in the lack of intra-seasonal variability, etc) and to
290 inter-model differences in simulations of the current climate and in climate change projections (especially changes in regional
291 precipitation and extreme events). More generally, a better understanding of how clouds couple to the circulation is expected
292 to improve our ability to answer the four science questions raised by the WCRP Grand Challenge on Clouds, Circulation and
293 Climate Sensitivity (Bony et al., 2015).

294 These questions provided the scientific motivation for the Clouds On/Off Klima Intercomparison Experiment (COOKIE)
295 project proposed by the European consortium EUCLIPSE and CFMIP (Stevens et al., 2012). The COOKIE experiments,
296 which have been run by four to eight climate models (depending on the experiment), switched off the cloud-radiative effects
297 (clouds seen by the radiation code -and the radiation code only- were artificially made transparent) in an atmospheric model
298 forced by prescribed SSTs. By doing so, the atmospheric circulation could feel the lack of cloud-radiative heating within the
299 atmosphere, but the land surface could also feel the lack of cloud shading, which led to changes in land surface temperatures
300 and land-sea contrasts. The change in circulation between On and Off experiments resulted from both effects, obscuring to
301 some degree the mechanisms through which the atmospheric cloud-radiative effects interact with the circulation for given
302 surface boundary conditions. As the longwave cloud-radiative effects are felt mostly within the troposphere (representing
303 most of the net atmospheric cloud-radiative heating) while the shortwave effects are felt mostly at the surface (e.g. L'Ecuyer
304 and McGarragh 2010; Haynes et al., 2013), we could better isolate the role of tropospheric cloud-radiative effects on the

305 circulation by running atmosphere-only experiments in which clouds are made transparent to radiation only in the longwave.
306 In this configuration, the models will have a shortwave cloud feedback but no longwave cloud feedback. We note that the
307 presence of clouds does affect the shortwave radiative heating of the atmosphere, although this is a much smaller effect than
308 its longwave equivalent (e.g. Pendergrass and Hartmann, 2014).

309 Therefore we propose in Tier 2 a set of simple experiments similar to the *amip*, *amip-p4K*, *aqua-control* and *aqua-p4K*
310 experiments within Tier 1, but in which cloud-radiative effects are switched off in the longwave part of the radiation code
311 while retaining those in the shortwave (Fermepin and Bony, 2014). Care should also be taken to remove the effects of cloud
312 on any longwave cooling used in other model schemes (e.g. turbulent mixing) if these are calculated independently of the
313 radiation scheme. These experiments will be referred to as *amip-lwoff*, *amip-p4K-lwoff*, *aqua-control-lwoff* and *aqua-p4K-*
314 *lwoff*. The analysis of idealized (aqua-planet) experiments will allow us to assess the robustness of the impacts found in more
315 realistic (AMIP) configurations. It will also facilitate the interpretation of the results using simple dynamical models or
316 theories, in collaboration with large-scale dynamicists (e.g. DynVar). The comparison of the inter-model spread of
317 simulations between the standard and ‘lwoff’ experiments for present-day and warmer climates will help to identify which
318 aspects of the inter-model spread depend on the representation of cloud-radiative effects, and which aspects do not, thus
319 better highlighting other sources of spread. An alternative method (proposed by Aiko Voigt) was also considered, in which
320 clear-sky heating rates would be applied in the atmosphere while retaining the all-sky fluxes at the surface. Although this
321 approach would potentially isolate the effects of cloud heating in the atmosphere more cleanly than the lwoff experiments
322 proposed here, it is yet to be demonstrated in a pilot study, and is considered more technically difficult to implement than the
323 lwoff experiments, which are very similar to those piloted by Fermepin and Bony, 2014.

324 **2.4 Abrupt +/-4% Solar Forced AOGCM experiments (Tier 2)**

325 Lead coordinators: Chris Bretherton, Roger Marchand, Bjorn Stevens

326
327 Science Question: How do responses in the climate system due to changes in solar forcing differ from changes due to CO₂,
328 and is the response sensitive to the sign of the solar forcing?
329

330 While rapid adjustments in clouds and precipitation can easily be separated from conventional feedbacks in SST forced
331 experiments, such a separation in coupled models is complicated by various issues, including the response of the ocean on
332 decadal timescales. A number of studies have examined cloud feedbacks in coupled models subject to a solar forcing, which
333 is generally associated with much smaller global cloud and precipitation adjustment, due to a smaller atmospheric absorption
334 for a given top of atmosphere forcing (e.g. Lambert and Faull, 2007; Andrews et al., 2010), but the regional cloud and
335 precipitation changes have yet to be rigorously investigated across models. Solar forcing also differs from greenhouse
336 forcing through its different fingerprint on the vertical structure of warming (Santer et al., 2013) and small changes in the
337 radiative heating near the tropopause may project measurably on tropospheric climate (e.g., Butler et al., 2010), for instance
338 by influencing the baroclinicity in the upper troposphere and thus the storm-tracks (Bony et al., 2015).

339 A +4% solar experiment *abrupt-solp4p* is proposed which is analogous to the *abrupt-4xCO2* experiment but rather than
340 changing CO₂ it would abruptly increase the solar constant by four percent and keep it fixed for 150 years, resulting in a
341 global mean radiative forcing of a similar magnitude to that due to CO₂ quadrupling. When changing the solar constant, the
342 shape of the spectral solar irradiance distribution should remain consistent with that in the piControl experiment. This
343 experiment complements the DECK *abrupt-4xCO2* experiment, tests the forcing feedback framework for analyzing climate
344 change, and would support our understanding of regional responses of the coupled system with and without CO₂ adjustments.
345 The complementary -4% abrupt solar forcing experiment (*abrupt-solm4p*) would allow the examination of feedback
346 asymmetry under climate cooling, and would also help with the interpretation of model responses to geo-engineering
347 scenarios and volcanic forcing, and of past climate signals.

348 **2.5 nonLinMIP abrupt 2xCO₂ and abrupt 0.5xCO₂ Experiments (Tier 2)**

349 Lead Coordinator: Peter Good

350
351 Science Question: To what extent is regional-scale climate change per CO₂ doubling state-dependent (nonlinear); what are
352 the associated mechanisms; and how does this affect our understanding of climate model uncertainty?
353

354 Recent studies with individual, or a small number of climate models, have found substantial nonlinearities in regional-scale
355 precipitation change (Good et al., 2012; Chadwick and Good, 2013), associated with robust physical mechanisms (Chadwick
356 and Good, 2013). Significant nonlinearity has also been found in global and regional-scale warming (e.g. Colman and
357 McAvaney, 2009; Jonko et al., 2013; Good et al., 2015; Meraner et al., 2013) and ocean heat uptake (Bouttes et al., 2015).

358 To address this science question we propose two new experiments for Tier 2, *abrupt 2xCO₂* and *abrupt 0.5xCO₂*. These are
359 the same as the DECK *abrupt4xCO2* experiment except that CO₂ concentrations are doubled and halved respectively relative
360 to the preindustrial control. These experiments are based on a proven analysis approach, including traceability of these
361 experiments to transient-forcing simulations (Good et al., 2016), to explore global and regional-scale nonlinear responses,
362 highlighting different behaviour under business-as-usual scenarios, mitigation scenarios and palaeoclimate simulations.
363 Additionally comparisons of the abrupt 2xCO₂ and abrupt 4xCO₂ experiments will help to establish the extent to which the

364 latter accurately estimates the equilibrium climate sensitivity to CO₂ doubling (e.g. Gregory et al., 2004, Block and
365 Mauritsen, 2013). Additional experiments (Good et al., 2016) may be proposed for Tier 2 in the future, or coordinated via
366 CFMIP outside of CMIP6. These include 100-year extensions to *abrupt-4xCO₂* and *abrupt-2xCO₂*; a 1% ramp-down from
367 the end of the *1pctCO₂* experiment; an abrupt step-down to 1xCO₂ from year 100 of the *abrupt-4xCO₂*. These would be used
368 to explore longer-timescale responses, quantify nonlinear mechanisms more precisely and understand the reversibility of
369 climate change.

370 **2.6 Feedbacks in AMIP experiments (Tier 2)**

371 Lead Coordinator: Timothy Andrews

372

373 Science question: Are climate feedbacks during the 20th century different to those acting on long term climate change?

374

375 Recent studies have shown significant time variation in climate feedbacks in response to CO₂ quadrupling (e.g. Andrews et
376 al., 2012; Geoffroy et al., 2013; Armour et al., 2013; Andrews et al., 2015). This raises the possibility that feedbacks during
377 the 20th century may be different to those acting on long term change, and hence has the potential to alleviate the apparent
378 discrepancy between estimates of climate sensitivity from comprehensive climate models and from simple climate models
379 fitted to observed warming trends (Collins et al., 2013). For example Gregory and Andrews, 2016 found that two models
380 forced with observed monthly 20th century SST and sea-ice variations simulated effective climate sensitivities of about 2K,
381 whereas these same models forced with patterns of long term SST change simulated effective climate sensitivities of over 3K
382 and 4K.

383 The previous CFMIP-2/CMIP5 design was unable to diagnose the time-variation of feedbacks of explicit relevance to the
384 historical period, because this requires the removal of the time varying forcing. To address this we propose an additional
385 experiment called '*amip-piForcing*' (amip pre-industrial forcing) following the design of Andrews 2014 and Gregory and
386 Andrews, 2016. This experiment is the same as the *amip* DECK experiment (i.e. using observed monthly updating SSTs and
387 sea-ice), but run for the period 1870-present and with constant pre-industrial forcings (i.e. all anthropogenic and natural
388 forcing boundary conditions identical to the *piControl* experiment). Since the forcing constituents do not change in this
389 experiment it readily allows a simple diagnosis of the simulated atmospheric feedbacks to observed SST and sea-ice changes,
390 which can then be compared to feedbacks representative of long term change and climate sensitivity (e.g. from *abrupt-4xCO₂*
391 or *amip-p4K*). The experiment has the additional benefit, by differencing with the standard *amip* run that includes time-
392 varying forcing agents, of providing detailed information on the transient effective radiative forcing and adjustments in
393 models during the AMIP period (Andrews, 2014). This can then be compared to the forcings diagnosed in the Radiative
394 Forcing Model Intercomparison Project (RFMIP, Pincus et al., 2016, who use a pre-industrial climate baseline) to test for any
395 dependence of forcing and adjustments on the climate state. Time-varying feedbacks in the *amip* experiment could
396 alternatively be diagnosed by subtracting a time-varying radiative forcing diagnosed from RFMIP experiments. However, the
397 *amip-piForcing* approach has the benefit of diagnosing the time-varying feedbacks over the full 1870-present period rather
398 than the last 36 years, and does so with reference to a single experiment, which reduces noise compared to that which would
399 be present with a double difference of the *amip* experiment and two RFMIP experiments. Also, the inclusion of CFMIP
400 process diagnostics in the *amip-piForcing* experiment will enable a deeper understanding of the factors underlying forcing
401 and feedback differences in the present and future climate.

402 **2.7 Time slice experiments for understanding regional climate responses to CO₂ (Tier 2)**

403 Lead Coordinators: Robin Chadwick, Hervé Douville and Christopher Skinner

404

405 Science questions:

- 406 • How do regional climate responses (e.g. of precipitation) in a coupled model arise from the combination of
407 responses to different aspects of CO₂ forcing and sea surface warming (uniform SST warming, patterned SST
408 warming, sea-ice change, direct CO₂ effect, plant physiological effect)?
- 409 • Which aspects of forcing/warming are most important for causing inter-model uncertainty in regional climate
410 projections?
- 411 • Can inter-model differences in regional projections be related to underlying structural or resolution differences
412 between models through improved process understanding, and could this help us to constrain the range of regional
413 projections?
- 414 • What impact do coupled model SST biases have on regional climate projections?

415

416 The CFMIP-2/CMIP5 set of idealised *amip* experiments (e.g. *amip4K*, *amipFuture*) have allowed the contribution of different
417 aspects of SST warming and increased CO₂ concentrations to the projections of fully coupled GCMs to be examined (e.g.
418 Bony et al., 2013; Chadwick et al., 2014; He and Soden, 2015). However the *amip* experiments were not designed to replicate
419 coupled GCM responses on a regional scale, and large discrepancies exist between the two in many regions, particularly
420 when individual models are examined instead of the ensemble mean (Chadwick, 2016). This is largely due to the choice of

421 present-day and future SST boundary conditions used in the amip experiments, as well as missing processes such as the plant
422 physiological response to CO₂, rather than the lack of air-sea coupling (Skinner et al., 2012).
423 We propose a new set of 7 30-year atmosphere-only time slice experiments, and one 36-year amip-style experiment, to
424 decompose the regional responses of each model's *abrupt-4xCO2* run into separate responses to each aspect of forcing and
425 warming (uniform SST warming, pattern SST change, sea-ice change, increased CO₂, plant physiological effect). These are
426 forced with monthly- and annually-varying monthly mean SSTs and sea ice, which reproduce regional precipitation patterns
427 more accurately than is possible using climatological SST forcing (Skinner et al., 2012). As well as allowing regional
428 responses in each individual model to be better understood, this set of experiments should prove especially useful for
429 understanding the causes of model uncertainty in regional climate change.

430 The experiments are:

431 1) *piSST* – An AGCM experiment with monthly- and annually-varying SSTs, sea-ice, atmospheric constituents and any
432 other necessary boundary conditions (e.g. vegetation if required) taken from a section of each model's own *piControl* run,
433 using the 30 years of *piControl* that are parallel to years 111-140 of its *abrupt-4xCO2* run. Note that dynamic vegetation (if
434 included in the model) should not be turned on in any of the *piSST* set of experiments;

435 2) *piSST-pxK* – same as *piSST*, but with a global spatially and temporally uniform SST anomaly applied on top of the
436 monthly- and annually- varying *piSST* SSTs. The magnitude of the uniform increase is taken from each model's global,
437 climatological annual mean open SST change between *abrupt-4xCO2* and *piControl* (using the mean of years 111-140 of
438 *abrupt-4xCO2*, and the parallel 30-year section of *piControl*). Sea-ice is unchanged from *piSST* values;

439 3) *piSST-4xCO2-rad* – same as *piSST* but CO₂ as seen by the radiation scheme is quadrupled;

440 4) *piSST-4xCO2* – same as *piSST* but with CO₂ quadrupled, and this increase is seen by both the radiation scheme and the
441 plant physiological effect. If a model does not include the plant physiological response to CO₂, then *piSST-4xCO2* can be
442 omitted from the set of *piSST* experiments for that model;

443 5) *a4SST* – same as *piSST*, but with monthly- and annually-varying SSTs taken from years 111-140 of each model's own
444 *abrupt-4xCO2* experiment instead of from *piControl* (sea ice is unchanged from *piSST*);

445 6) *a4SSTice* – same as *piSST*, but with monthly- and annually-varying SSTs and sea-ice taken from years 111-140 of each
446 model's own *abrupt-4xCO2* experiment instead of from *piControl*;

447 7) *a4SSTice-4xCO2* – same as *piSST*, but with monthly- and annually-varying SSTs and sea-ice taken from years 111-140
448 of each model's own *abrupt-4xCO2* experiment instead of from *piControl*. CO₂ is also quadrupled, and is seen by both the
449 radiation scheme and the plant physiological effect (if included in the model). *a4SSTice-4xCO2* is used to establish whether a
450 time slice experiment can adequately recreate the coupled *abrupt-4xCO2* response in each model, and then forms the basis for
451 a decomposition using the other experiments. The time slice experiments can be combined in various ways to isolate the
452 climate response to each individual aspect of forcing and warming. For example the response to SST pattern change is given
453 by taking the difference between *a4SST* and *piSST-pxK*, and the plant physiological response is found by taking the
454 difference between *piSST-4xCO2* and *piSST-4xCO2-rad*.

455 8) We also propose an additional amip based experiment, *amip-a4SST-4xCO2*: the same as amip, but a patterned SST
456 anomaly is applied on top of the monthly- and annually-varying amip SSTs. This anomaly is a monthly climatology, taken
457 from each model's own *abrupt-4xCO2* run minus *piControl* (using the mean of years 111-140 of *abrupt-4xCO2*, and the
458 parallel 30-year section of *piControl*). CO₂ is quadrupled, and the increase in CO₂ is seen by both the radiation scheme and
459 vegetation. Comparison of *amip-a4SST-4xCO2* and *a4SSTice-4xCO2* should help to illuminate the impact of SST biases on
460 regional climate responses in each model, and how this contributes to inter-model uncertainty.

461 3 CFMIP Recommended Diagnostic Outputs for CMIP experiments

462 The CFMIP-3 specific diagnostic request is designed to address the following questions: 1) How well do clouds and other
463 relevant variables simulated by models agree with observations? 2) What physical processes and mechanisms are important
464 for a credible simulation of clouds, cloud feedbacks and cloud adjustments in climate models? 4) Which models have the
465 most credible representations of processes relevant to the simulation of clouds? 5) How do clouds and their changes interact
466 with other elements of the climate system?

467 The set of diagnostic outputs recommended for CFMIP-3 is based on that from CFMIP-2, with some modifications. The
468 request outlined below is in three parts. The first part describes an updated set of CFMIP process diagnostics (based on those
469 in CFMIP-2 which are documented at http://cmip-pcmdi.llnl.gov/cmip5/output_req.html) in terms of the various groups of
470 variables and the experiments in which they are requested. This set was drawn up by the CFMIP committee and ratified by
471 the modelling groups following a presentation at the 2014 CFMIP meeting. The second part describes recommendations for
472 COSP outputs in the CFMIP-3, CMIP DECK and CMIP6 Historical experiments. The third part describes additional
473 diagnostics requested for evaluation of mean diurnal cycle of tropical clouds and radiation. The summaries below give an
474 overview of the diagnostic request; however the definitive and detailed specification is documented in the CMIP6 data
475 request, available at <https://www.earthsystemcog.org/projects/wip/CMIP6DataRequest> (Juckes et al., in preparation.) The
476 changes in the CFMIP-3 diagnostics relative to those requested for CFMIP-2 are additionally motivated and detailed in the
477 CFMIP CMIP6 proposal document which is available from the CFMIP website.

478 CMIP mandates that for participation in the CFMIP-3, modelling groups must commit to performing all of the Tier 1
479 experiments. In recognition that sufficient resources are not available for all groups to prepare all of the CFMIP-3 specific
480 diagnostics, these diagnostics are considered to be Tier 2, i.e., not compulsory for participation in CFMIP-3. Nonetheless,
481 these diagnostics are extremely valuable and all groups with the capacity to do so are very strongly encouraged to provide the
482 additionally requested CFMIP-3 specific diagnostics.

483

484 In the case where CFMIP-3 specific outputs are requested in DECK and CMIP6-Historical experiments, and modelling
485 groups run more than one ensemble member of an experiment, we request that each set of CFMIP-3 specific outputs are
486 submitted for one ensemble member only. Having different CFMIP variables in different ensemble members is acceptable,
487 but submitting them all in the same ensemble member is preferable. We request that the modelling groups provide
488 information on which CFMIP diagnostic sets are submitted in which ensemble members so that this information can be made
489 available to those who may be analyzing the output. Our analysis plans for the CFMIP diagnostic outputs in the CMIP
490 DECK, CMIP6 Historical and CFMIP-3 experiments, including details of the CFMIP Diagnostics Code Catalogue are
491 summarised in Appendix A.

492 3.1 Process outputs

493 In CFMIP-2, instantaneous high frequency ‘cfSites’ outputs were requested for 120 locations in the *amip*, *amip4K*,
494 *amipFuture* and *amip4xCO2* experiments, and for 73 locations along the Greenwich meridian in the aquaplanet experiments,
495 to support understanding and evaluation of clouds and their interactions with convection and other processes. The 120
496 locations include the locations of instrumented sites (ARM and CloudNet stations, Dome C, etc), the transect associated with
497 the GCSS Pacific Cross-section Intercomparison (GPCI), past field campaigns (DYCOMS-II, NARVAL, HOPE, VOCALS,
498 ASTEX and AMMA transects, TOGA-COARE, RICO, etc) and a number of climate regimes that contribute substantially to
499 the inter-model spread of cloud feedbacks in climate change (Webb et al., 2015a). These outputs have so far been used to
500 evaluate the models with in-situ measurements (e.g. Nuijens et al., 2015a, Nuijens et al., 2015b, Neggers et al., 2015), to
501 investigate the diurnal cycle of cloud feedbacks (Webb et al., 2015a) and to compare cloud feedbacks in climate models with
502 SCM and LES outputs from CGILS (Dal Gesso et al., 2015). We have added St. Helena to the list of locations in light of
503 upcoming field work, increasing the total number of locations to 121 for CFMIP-3. A text file containing the list of locations
504 is available in the Supplementary Information and on the CFMIP website; these are also presented graphically in Figure 2.

505 For CFMIP-3 cfSites outputs are now requested for one ensemble member of the *amip* DECK experiment, and the *amip*-
506 *p4K* and *amip-4xCO2* experiments. Outputs should be provided for the full duration of each experiment. The sampling
507 interval should be the integer multiple of the model time step that is nearest to 30 minutes and divides into 60 minutes with
508 no remainder: e.g. 30 minutes for a 30, 15 or 10 minute time step or 20 minutes for a 20 minute time step. Outputs should be
509 instantaneous (i.e. not time means) and from nearest grid box (i.e. no spatial interpolation). We have dispensed with the
510 cfSites outputs in the aquaplanet and *amip-future4K* experiments because these have been less widely used compared to those
511 from the other experiments.

512 The cfSites outputs from CFMIP-3 provide instantaneous outputs of a range of quantities (including temperature and
513 humidity tendency terms) in experiments which can be used to evaluate the present day relationships of clouds to cloud
514 controlling factors using in situ measurements, and at the same time explore how these relationships affect cloud feedbacks
515 and cloud adjustments. An increasing wealth of observational data with which to evaluate the models using these outputs is
516 available or in the planning stage, for example from the Barbados Cloud Observatory (Stevens et al., 2015) the ARM
517 Program (e.g. Wood et al., 2015; Marchand et al., 2015) or within the German national project on high-definition clouds and
518 precipitation for climate-prediction, HD(CP)², inclusive of its observational prototype experiment (HOPE), and which has
519 collected observations over Germany following conventions adopted for CMIP (Andrea Lammert, personal communication).

520 CFMIP-2 also requested cloud, temperature and humidity tendency terms from convection, radiation, dynamics etc. in the
521 *amip*, *amip4K*, *amipFuture* and *amip4xCO2*, *aquaControl*, *aqua4xCO2* and *aqua4K* experiments, as global monthly mean
522 outputs and high frequency outputs at fixed locations (Bony et al., 2011). Upward and downward radiative fluxes on model
523 levels were also requested in these experiments, and for instantaneous CO₂ quadrupling in the *amip* experiment only.
524 Temperature and humidity tendency terms in particular have been shown to be useful for understanding the roles of different
525 parts of the model physics in cloud feedbacks (e.g. Webb and Lock 2013; Demoto et al., 2013; Sherwood et al., 2014; Brient
526 et al., 2015) and cloud adjustments (e.g. Kamae and Watanabe 2012; Ogura et al., 2014) as well as in understanding clouds
527 and circulation in the present climate (e.g. Williams et al., 2013; Oueslati and Bellon, 2013; Xavier et al., 2015). They have
528 also been used to understand regional warming patterns such as polar amplification in coupled models (e.g. Yoshimori et al.,
529 2014).

530 In CFMIP-3 we have improved the definitions of the temperature and humidity tendency terms, and added some additional
531 terms such as clear-sky radiative heating rates to more precisely quantify the contributions of different processes to the
532 temperature and humidity budget changes underlying cloud feedbacks and adjustments. We have dispensed with the cloud
533 water tendency terms because these have been less widely used than the temperature and humidity tendencies.

534 A shortcoming of the CMIP5 protocol was that we were unable to interpret the physical feedback mechanisms in coupled
535 model experiments due to a lack of process diagnostics. For this reason in CMIP6 we are requesting these budget terms in
536 the DECK *abrupt-4xCO2* experiment and the pre-industrial control as well as one ensemble member of the *amip* DECK
537 experiment, and all of the CFMIP-3 experiments listed in Sections 2.1-2.6.

538 Clustering approaches (e.g., Jakob and Tselioudis, 2003) are now commonly used for assessing the contributions of
539 different cloud regimes (e.g. stratocumulus, trade cumulus, frontal clouds, etc) to present day biases in cloud simulations and
540 to inter-model differences in cloud feedbacks (e.g. Williams and Webb 2009, Tsushima et al., 2013, Tsushima et al., 2015).
541 We have also added some additional daily 2D fields to the standard package of CFMIP daily outputs to allow further
542 investigation of feedbacks between clouds and aerosols associated with the changing hydrological cycle (aerosol loadings and
543 cloud top effective radii/number concentrations) and a clearer diagnosis of the roles of convective and stratiform clouds
544 (convective vs. stratiform ice and condensed water paths and cloud top effective radii/number concentrations).

546 This section motivates and summarizes the COSP outputs requested from the DECK, and CMIP6 historical and CFMIP-3
 547 experiments as well as a corresponding set of observations.

548 There is no unique definition of clouds or cloud types, neither in models nor in observations. Therefore, to compare models
 549 with observations, and even to compare models with each other, it is necessary to use a consistent definition of clouds
 550 between the model and the satellite product in question (i.e., be “definition-aware”). Further complicating matters - climate
 551 model grid boxes (typically 1 degree) are much larger than the scales over which many satellite observations are made
 552 (typically <10 km). As a result, one must downscale the climate model cloud properties to the observation scale (i.e., be
 553 “scale-aware”). The CFMIP Observation Simulator Package (COSP) enables definition-aware and scale-aware comparisons
 554 between models and multiple sets of observations by producing cloud diagnostics from model simulations that are
 555 quantitatively comparable to a variety of satellite products from ISCCP, CloudSat, CALIPSO, MODIS, MISR and Parasol
 556 (Bodas-Salcedo et al., 2011). COSP enables a more quantitative comparison of model outputs with satellite cloud products,
 557 which often sub-sample low level clouds in the presence of high level clouds due to the effects of cloud overlap and
 558 attenuation (e.g. Yu et al., 1996). COSP also provides histograms of various cloud properties as a function of height or
 559 pressure which are directly comparable with satellite products and cannot be calculated correctly from time mean model
 560 outputs. The multiple simulators within COSP allow a multi-faceted evaluation of clouds in models whereby the strengths
 561 and weaknesses of different satellite products may be considered together.

562 COSP is increasingly being used not only for model intercomparison activities but as part of the model development and
 563 evaluation process by modelling groups (e.g. Marchand et al., 2009; Zhang et al., 2010; Kay et al., 2012; Franklin et al.,
 564 2013; Lacagnina and Selten, 2014; Nam et al., 2014; Williams et al., 2015, Konsta et al., 2015). Many of the standard
 565 monthly and daily COSP outputs have been shown to be valuable in the CMIP5 experiments, not only for cloud evaluation,
 566 allowing a detailed evaluation of clouds and precipitation, and their interaction with radiation (e.g. Nam et al., 2012; Cesana
 567 and Chepfer, 2012; Kay et al. 2012; Klein et al., 2013; Tsushima et al., 2013; Gordon and Klein, 2014; Lin et al., 2014;
 568 Bodas-Salcedo et al., 2014; Bellomo and Clement, 2015), but also in quantifying the contributions of different cloud types to
 569 cloud feedbacks and forcing adjustments in climate change experiments (e.g. Zelinka et al., 2013; Zelinka et al., 2014;
 570 Chepfer et al., 2014; Tsushima et al., 2015). For a full list of studies that use COSP diagnostics for model evaluation and
 571 feedback analysis please refer to the ‘CFMIP publications’ section of the CFMIP website.

572 Here we will give only a brief overview of the COSP request; readers interested in the complete details of the data request
 573 are referred to the Earth System CoG website (<https://www.earthsystemcog.org/projects/wip/CMIP6DataRequest>). The
 574 COSP data request for the CMIP DECK and CMIP6 has been designed to span model evaluation across different space and
 575 time scales. Monthly-mean diagnostics allow for the evaluation and intercomparison of large-scale distributions of cloud
 576 properties and their interaction with radiation. High-frequency model outputs (daily, 3-hourly) are aimed at a process-oriented
 577 evaluation (e.g. Bodas-Salcedo et al., 2012) and offer the opportunity of exploiting the synergy between multiple instruments
 578 (e.g. Konsta et al., 2015). Recent observational developments have improved our capability to retrieve cloud radiative
 579 properties. In particular, new methodologies for cloud phase identification are available for CALIPSO and MODIS, and
 580 COSP has been enhanced to provide diagnostics that are compatible with these new observational datasets (Cesana and
 581 Chepfer, 2013). These new diagnostics will help elucidate some open questions regarding the role of cloud phase in model
 582 biases (Ceppi et al., 2016; Bodas-Salcedo et al., in press).

583
 584 Within CFMIP-3 COSP output is requested from six simulators as follows:

- 585 • ISCCP: pseudo-retrievals of cloud top pressure (CTP) and cloud optical thickness (tau) (Klein and Jakob 1999;
 586 Webb et al., 2001).
- 587 • CloudSat: a forward model for radar reflectivity as a function of height (Haynes et al., 2007).
- 588 • CALIPSO (Chepfer et al., 2008; Cesana and Chepfer, 2013): forward model for lidar scattering ratio as function of
 589 height, and cloud phase retrieval.
- 590 • MODIS: pseudo-retrievals of CTP, effective particle size and tau as function of phase (Pincus et al., 2012).
- 591 • MISR: pseudo-retrievals of cloud top height (CTH) and tau (Marchand and Ackerman, 2010).
- 592 • PARASOL: simple forward model of mono-directional reflectance (Konsta et al., 2015).

593
 594 The main difference to CFMIP-2 is that output is requested from a greater number of simulators and longer periods of
 595 simulated time. MISR provides more accurate retrievals of cloud-top-height for low-level and mid-level clouds, and more
 596 reliable discrimination of mid-level clouds from other clouds, while MODIS provides better retrievals of high-level clouds.
 597 ISCCP and MISR histograms can be combined to separate optically-thin high-level clouds into multi-layer and single-layer
 598 categories (Marchand et al. 2010). Aerosol schemes are becoming more complex, with more elaborate representations of
 599 cloud-aerosol interactions. This makes the evaluation of the phase partitioning an important aspect of model evaluation, and
 600 height-resolved partitioning estimates from the CALIPSO simulator are included in the COSP request. Cloud phase and
 601 particle size estimates from the MODIS simulator were not available in CFMIP-2 but may prove a useful complement to
 602 investigate cloud-aerosol interactions by virtue of greater geographic sampling and longer time records. Many of the COSP
 603 diagnostics are now requested for the entire lengths of the DECK, CMIP6 Historical and CFMIP-3 experiments to support the
 604 quantification and interpretation of cloud feedbacks and cloud adjustments in a broader context. The new inclusion in this
 605 COSP request of a long time series of three-dimensional cloud fractions will facilitate the comparison of cloud trends with the
 606 observational record (Chepfer et al., 2014). More details of all the changes with respect to CFMIP-2 can be found in the

607 proposal of the CMIP6-Endorsed MIPs, available from the CMIP6 website (<http://www.wcrp-climate.org/wgcm-cmip/wgcm-cmip6>).

608
609
610 The COSP output is in six variable groups:

- 611 1) *cfMon_sim*: monthly means of ISCCP 2D diagnostics (cloud fraction, cloud albedo, and cloud top pressure), ISCCP CTP-tau histogram, and CALIPSO 2D and 3D cloud fractions.
- 612 2) *cfDay_2d*: daily means of ISCCP and CALIPSO 2D diagnostics, and PARASOL reflectances.
- 613 3) *cfDay_3d*: daily means of ISCCP and CALIPSO 3D diagnostics.
- 614 4) *cfMonExtra*: monthly means of CloudSat reflectivity and CALIPSO scattering ratio histograms as function of height, CALIPSO 3D cloud fractions by phase, MODIS 2D cloud fractions, MODIS CTP-tau histogram and size-tau histograms by phase, MISR CTH-tau histograms, and PARASOL reflectances.
- 615 5) *cfDayExtra*: daily means of CALIPSO total cloud fraction, MODIS CTP-tau histogram and size-tau histograms by phase, and PARASOL reflectances.
- 616 6) *cf3hrSim*: 3-hourly instantaneous diagnostics of ISCCP CTP-tau histograms, MISR CTH-tau histograms, MODIS CTP-tau histogram and size-tau histograms by phase, CALIPSO 2D and 3D cloud fractions, CloudSat reflectivity and CALIPSO scattering ratio histograms as function of height, and PARASOL reflectances.

625 The variable groups *cfMon_sim* and *cfDay_2d* are requested for all years in the *amip* experiment performed as part of the DECK and the CMIP6-Historical experiments, and for 140 years of the *piControl*, *1pctCO2*, and *abrupt-4xCO2*. These are requested for one ensemble member only from these experiments. They are also requested in all of the CFMIP experiments listed in Sections 2.1-2.6 above. *cfDay_3d* is requested in one ensemble member of the DECK *amip* experiment and in the CFMIP *amip-p4K* and *amip-4xCO2* experiments. *cfMonExtra* and *cfDayExtra* are requested for all years of one ensemble member of the *amip* DECK experiment, and *cf3hrSim* for the year 2008 only. (Please note that in the full data request these variable groups are in many cases split into a number of sub-tables. As noted above, the formal data request provides the definitive specification of the model outputs.)

633 COSP is available via the CFMIP website (<https://www.earthsystemcog.org/projects/cfmip>). Version 1.4 is a stable code release that was made available well in advance of CMIP6 at the request of the modelling groups. Small updates are required to enable some new diagnostics requested by CFMIP3/CMIP6, most notably joint histograms of particle size and optical thickness from the MODIS simulator; with these updates the code is known as version 1.4.1. Modeling centers are encouraged to update to COSP 1.4.1 to provide these new diagnostics but may provide results from COSP 1.4.

638 Developed over the last few years, COSP 2 substantially revises the infrastructure for integrating satellite simulators in climate models. COSP 2 makes many fewer inherent assumptions about the model representation of clouds than do previous versions but contains an optional interface allowing it to be used as a drop-in replacement for COSP 1.4 or COSP 1.4.1. At the time of this writing COSP 2 is undergoing final testing in two climate models. Availability of the final version will be announced on the CFMIP website and modelling groups are free to adopt it for use in CFMIP at that time.

643 The CFMIP community has developed a set of observational datasets available via the CFMIP-OBS website (<http://climserv.ipsl.polytechnique.fr/cfmip-obs/>) that are defined consistently with the COSP diagnostics and the CFMIP data request in terms of vertical grids and time averaging periods. These are mostly reported as monthly means although some are reported at higher temporal resolution for process oriented model evaluations (e.g. Konsta et al., 2012). Table 3 summarizes the datasets relevant to the COSP CMIP6 data request. Some of the CFMIP-OBS datasets listed in Table 3 (CALIPSO, CloudSat, ISCCP, PARASOL) are also available from the ESGF as part of the obs4MIPs project (Teixeira et al., 2014). These datasets are periodically updated to include more recent data from the relevant satellites, many of which are still operational. Please refer to the CFMIP-OBS website for updates.

652 3.3 Monthly Mean Diurnal Cycle Outputs

653 Climate models have difficulties representing the diurnal cycle of convective clouds over land (Yang and Slingo, 2001; Stratton and Stirling, 2011), but its evaluation is not possible with sun-synchronous satellites. Geostationary satellites provide high-frequency sampling that can be used to evaluate model biases in the diurnal cycle of clouds and radiation (albeit over a limited area). The Geostationary Earth Radiation Budget instrument (GERB; Harries et al., 2005) measures the top of atmosphere (TOA) radiation budget from a geostationary orbit at 0E at 15 minute frequency, which provides a unique view of tropical convection over Africa. The variable group *cf1hrClimMon* requests monthly mean diurnal cycle of TOA radiative fluxes (all-sky and clear sky) for the entire length of the *amip* DECK experiment. The radiative fluxes are hourly UTC means. The ‘average day’ for each month of the simulation is then constructed by averaging each UTC hourly mean over the entire month. These diagnostics will be directly comparable with GERB measurements.

664 4. Summary

665 The primary goal of CFMIP is to inform improved assessments of cloud feedbacks on climate change. This involves bringing climate modelling, observational and process modelling communities closer together and providing better tools and community support for understanding and evaluation of clouds and cloud feedbacks simulated by climate models. CFMIP

668 supports ongoing coordinated model inter-comparison activities by recommending experiments and model output diagnostics
669 for CMIP, designed to support the understanding and evaluation of cloud processes and cloud feedbacks in models. The
670 CFMIP approach is also increasingly being used to understand other aspects of climate change, and so a second objective has
671 now been introduced, to improve understanding of circulation, regional-scale precipitation, and non-linear changes. CFMIP
672 proposes a number of experiments and model outputs for CMIP6, building on and extending those which were part of
673 CMIP5.

674 A compact set of Tier 1 experiments are proposed address the question: “1) What are the physical mechanisms underlying
675 the range of cloud feedbacks and cloud adjustments predicted by climate models, and which models have the most credible
676 cloud feedbacks?” The Tier 1 experiments (*amip-p4K*, *amip-4xCO2*, *amip-future4K*, *aqua-control*, *aqua-4xCO2* and *aqua-
677 p4K*) retain the idealized experimental hierarchy of the CFMIP-2/CMIP5 experiments while building on the DECK AMIP
678 experiment. A number of Tier 2 experiments are proposed to address additional science questions. An amip uniform minus
679 4K experiment is proposed to address the question “2) Are cloud feedbacks consistent for climate cooling and warming, and
680 if not, why?” Atmosphere-only experiments with clouds made transparent to longwave radiation address the question “3)
681 How do cloud-radiative effects impact the structure, the strength and the variability of the general atmospheric circulation in
682 present and future climates?” Abrupt +/-4% Solar Forced AOGCM experiments are proposed for the question “4) How do
683 responses in the climate system due to changes in solar forcing differ from changes due to CO₂, and is the response sensitive
684 to the sign of the solar forcing?” abrupt 2xCO₂ and abrupt 0.5xCO₂ experiments are proposed to address the question “5) To
685 what extent is regional-scale climate change per CO₂ doubling state-dependent (nonlinear), and why?” Other experiments and
686 questions proposed include: AMIP with preindustrial forcing “6) Are climate feedbacks during the 20th century different to
687 those acting on long term climate change and climate sensitivity?”; Time slice experiments forced with SSTs from
688 preindustrial and *abrupt-4xCO2* simulations “7) How do regional climate responses (of e.g. precipitation) in a coupled model
689 arise from the combination of responses to different aspects of CO₂ forcing and warming (uniform SST warming, pattern SST
690 warming, direct CO₂ effect, plant physiological effect, sea-ice change)?”

691 The CFMIP experiments in CMIP6 will continue to include outputs from the CFMIP Observational Simulator Package
692 (COSP) to support robust scale-aware and definition-aware evaluation of modelled clouds with observations and to relate
693 cloud feedbacks to observed quantities. COSP outputs are also proposed for inclusion in the DECK and CMIP6 Historical
694 experiments. Process diagnostics including ‘cfSites’ high frequency outputs at selected locations and temperature and
695 humidity budget terms from radiation, convection, dynamics, etc. are also retained from CMIP5. These will help to address
696 the following questions: 1) How well do clouds and other relevant variables simulated by models agree with observations?
697 2) What physical processes and mechanisms are important for a credible simulation of clouds, cloud feedbacks and cloud
698 adjustments in climate models? 4) Which models have the most credible representations of processes relevant to the
699 simulation of clouds? 5) How do clouds and their changes interact with other elements of the climate system?
700 By continuing the CFMIP experiments and diagnostic outputs within CMIP6 we hope to apply the well established aspects of
701 the CFMIP approach to a larger number of climate models. Additionally we have proposed new experiments to investigate a
702 broader range of questions relating to the Grand Challenge on Clouds, Circulation and Climate Sensitivity. We hope that the
703 modelling community will participate fully in CFMIP via CMIP6 so as to maximize the relevance of our findings to future
704 assessments of climate change.

705 Code and Data Availability

706 COSP is published under an open source license via GitHub (please see the CFMIP website for details). The model output
707 from the DECK, CMIP6 historical and CFMIP-3 simulations described in this paper will be distributed through the Earth
708 System Grid Federation (ESGF) with digital object identifiers (DOIs) assigned. As in CMIP5, the model output will be freely
709 accessible through data portals after registration. In order to document CMIP6's scientific impact and enable ongoing support
710 of CMIP, users are obligated to acknowledge CMIP6, the participating modelling groups, and the ESGF centres (see details
711 on the CMIP Panel website at <http://www.wcrp-climate.org/index.php/wgcm-cmip/about-cmip>). Further information about
712 the infrastructure supporting CMIP6, the metadata describing the model output, and the terms governing its use are provided
713 by the WGCM Infrastructure Panel (WIP) in their invited contribution to this Special Issue. Along with the data itself, the
714 provenance of the data will be recorded, and DOIs will be assigned to collections of output so that they can be appropriately
715 cited. This information will be made readily available so that published research results can be verified and credit can be
716 given to the modelling groups providing the data. The WIP is coordinating and encouraging the development of the
717 infrastructure needed to archive and deliver this information. In order to run the experiments, datasets for natural and
718 anthropogenic forcings are required. These forcing datasets are described in separate invited contributions to this Special
719 Issue. The forcing datasets will be made available through the ESGF with version control and DOIs assigned.

720
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734 **Appendix A: Analysis Plan and CFMIP Diagnostic Codes Catalogue**

735 CFMIP-2 analysis activities are ongoing and the CFMIP community is ready to analyse CFMIP-3 data at any time. We would
736 like modelling groups to perform the proposed CFMIP-3 experiments at the same time or shortly after their DECK and
737 CMIP6 Historical experiments. Subsequent CFMIP experiments which are not included in CMIP6 will build on the proposed
738 DECK and CMIP6/CFMIP experiments and some will start as soon as CMIP6 DECK experiments start to become available.
739 We envisage a succession of CFMIP related intercomparisons addressing different questions arising from the Grand
740 Challenge spanning the duration of CMIP6.

741 We plan to scientifically analyze, evaluate and exploit the proposed experiments and diagnostic outputs, and have
742 identified lead coordinators within CFMIP for different aspects of this activity. The lead coordinators are responsible for
743 encouraging analysis of the relevant experiments as broadly as possible across the scientific community. While they may lead
744 some analysis themselves, they do not have any first claim on analysing or publishing the results. All interested investigators
745 are encouraged to exploit the data from these experiments. While investigators may wish to liaise with the lead coordinators
746 to avoid duplicating work that others are doing, this is not a requirement. An overview of the proposed evaluation/analysis of
747 the CMIP DECK, CMIP6 Historical and CFMIP CMIP6 experiments follows:

748 CFMIP will continue to exploit the CMIP DECK and CMIP6 experiments to understand and evaluate cloud processes and
749 cloud feedbacks in climate models. The wide range of analysis activities described above in the context of CFMIP-2 will be
750 continued in CFMIP-3 using the CMIP DECK and CMIP6 experiments, allowing the techniques developed in CFMIP-2 to
751 applied to an expanding number of models, including the new generation of models currently under development. These
752 activities will include evaluation of clouds using additional simulators, investigation of cloud processes and cloud
753 feedback/adjustment mechanisms using process outputs (cfSites, tendency terms, etc). The inclusion of COSP and budget
754 tendency terms in additional DECK experiments (e.g. *abrupt-4xCO2*) will enable the CFMIP approach to be applied to a
755 wider range of experimental configurations. Lead coordinator: Mark Webb.

756 Analysis of the +/-4% solar forcing runs will include an evaluation of both rapid adjustments and longer-term responses on
757 global and regional top-of-atmosphere radiative fluxes, cloud types (using ISCCP and other COSP simulators) and
758 precipitation characteristics, as well as comparison of these responses with responses in DECK *abrupt-4xCO2* experiments.
759 GeoMIP and SolarMIP have expressed a strong interest in these CFMIP experiments and joint analysis of these CFMIP
760 experiments with GeoMIP and SolarMIP experiments is anticipated, specifically with the goal of determining to what degree
761 results from abrupt solar forcing only experiments and abrupt CO₂ only experiments can be used to predict what happens
762 when both forcing are applied simultaneously, as done in the GeoMIP experiments. Lead coordinators: Chris Bretherton,
763 Roger Marchand and Bjorn Stevens.

764 Analysis of nonlinear climate processes is discussed in detail by Good et al., 2016. This includes a method for validating
765 traceability of abrupt CO₂ experiments to transient simulations, which is also recommended as a standard test of the DECK
766 *abrupt-4xCO2* experiment. Analysis will primarily involve comparing the *abrupt-4xCO2*, *abrupt-2xCO2* and *abrupt-*
767 *0p5xCO2* experiments over the same timescale. Lead coordinator: Peter Good.

768 Analysis of *amip-piForcing* has already been performed in detail for two models in Andrews, 2014 and Gregory and
769 Andrews (submitted). We propose to use this as a starting point for a multi-model analysis. Lead coordinator: Timothy
770 Andrews.

771 An overview analysis of regional responses and model uncertainty in the piSST set of experiments will be carried out by
772 the coordinators, in collaboration with members of contributing modelling groups. We anticipate that further detailed analysis
773 on the processes at work in different regions will be carried out by a variety of research groups with interest and expertise in a
774 particular region: for example a set of similar experiments has previously been used to examine the climate response of the
775 West African monsoon in CCSM3 (Skinner et al., 2012). The piSST set of experiments have already been successfully run
776 using the Met Office, NCAR and CNRM CMIP5 models. Lead Coordinators: Robin Chadwick, Hervé Douville and
777 Christopher Skinner.

778 The analysis of the COOKIE experiments will be reviewed by the coordinators in collaboration with members of the
779 contributing modelling groups. The role of longwave atmospheric cloud-radiative effects in large-scale circulations, regional
780 precipitation patterns and the organisation of tropical convection will be investigated in the current climate and in climate
781 change, with the aim of highlighting both robust effects and sources of uncertainties in the model responses. Lead
782 coordinators: Sandrine Bony and Bjorn Stevens.

783 When analyzed together with the *amip-p4K* experiment, the *amip-m4K* experiment allows the CFMIP process diagnostics
784 to be used to understand for asymmetries in the climate response to warming and cooling which have been noted in PMIP
785 experiments. These might arise from cloud phase responses in middle- and high-latitude clouds or from the adiabatic cloud
786 liquid water path response feedback which is important over land regions and which would be expected to be weaker with
787 cooling because of the non-linearity in the Clausius-Clapeyron relation. Lead coordinators: Mark Webb and Bjorn Stevens.

788 The COSP data request for the *amip* DECK experiment will allow a comprehensive multi-model evaluation of clouds and
789 radiation, following on from CMIP5 studies (e.g. Klein et al., 2013; Bodas-Salcedo et al., 2014). The COSP data request for
790 the other experiments (e.g. *amip-p4K*, *abrupt-4xCO2*, etc.) permits evaluation of cloud feedbacks and adjustments by cloud

791 type (Zelinka et al., 2013, Tsushima et al., 2015) or cloud trends (Chepfer et al., 2014). New COSP diagnostics have been
 792 used in single-model analyses: cloud phase diagnostics (Cesana and Chepfer, 2013); MISR simulator outputs to evaluate
 793 cloud fraction and multilayer clouds (Marchand and Ackerman, 2010); CALIPSO vertical distribution of cloud fraction for
 794 the study of cloud trends (Chepfer et al., 2014). These studies will be used as starting points for multi-model analyses. The
 795 COSP Project Management Committee co-chairs will coordinate and encourage the exploitation of these resources. Lead
 796 coordinators: Alejandro Bodas-Salcedo and Steve Klein.

797 Analysis of output from CFMIP and CMIP6 experiments will also be facilitated by sharing of diagnostic codes via the
 798 CFMIP Diagnostics Code Catalogue (accessible via the CFMIP website <http://www.earthsystemcog.org/projects/cfmip/>).
 799 This is a catalogue of programs written by various members of the CFMIP community, implementing a number of diagnostic
 800 approaches from published studies. These include daily cloud clustering evaluation metrics based on ISCCP and ISCCP
 801 simulator outputs (Williams and Webb, 2009, Tsushima et al., 2013), error metrics for total cloud amount, longwave and
 802 shortwave cloud properties (Klein et al., 2013), process oriented evaluation of clouds using A-train instantaneous
 803 observations (Konsta et al., 2012), quality control and low-cloud diagnostics (Nam et al., 2012; Nam and Quaas, 2012),
 804 sensitivity of low cloud cover to estimated inversion strength and SST (Qu et al., 2014) and cloud radiative kernels (Zelinka
 805 et al., 2012). Any codes which implement diagnostics which are relevant to analysing clouds, circulation and climate
 806 sensitivity in models and which are documented in peer reviewed studies are eligible for inclusion in the catalogue, and we
 807 welcome additional contributions to further support community analysis of CMIP6 outputs.

808 APPENDIX B: Aquaplanet Experimental Design

809 Aquaplanets are Earth-like planets with completely water-covered surfaces. They are often used as idealized configurations
 810 of atmospheric GCMs, and in this context the usual convention is that landmasses and topography are removed. Although
 811 many flavours of aquaplanet configurations exist, another convention is to retain as much of the atmospheric model's
 812 formulation as possible. That is, the numerical grid, dynamical core, and parameterized physics are all used just as in realistic
 813 climate simulations.

814
 815 The Tier 1 aquaplanet experiments follow the same experimental design as CMIP5/CFMIP-2 (Medeiros et al., 2015). Those,
 816 in turn, were closely related to previous aquaplanet descriptions. In particular, the control configuration closely follows the
 817 AquaPlanet Experiment protocol (Blackburn and Hoskins, 2013) using a prescribed SST pattern described by Neale and
 818 Hoskins (2000). Two additional runs parallel the CFMIP-2 *amip4K* and *amip4xCO2* experiments: a uniform 4K warming and
 819 a quadrupling of atmospheric CO₂.

820 Here we provide the detailed experimental protocol for the three aquaplanet simulations that are part of Tier 1. We note
 821 again that these follow the APE protocol and CMIP5/CFMIP-2, and therefore largely mirror previous descriptions in
 822 Blackburn and Hoskins (2013), Williamson et al. (2012), and Medeiros et al. (2015).

823 Orbital parameters are set to perpetual equinox conditions. This is usually achieved by setting eccentricity and obliquity to
 824 zero to define a circular orbit and insolation independent of calendar. The diurnal cycle is retained. Insolation is based on a
 825 non-varying solar constant of 1365 W m⁻².

826 The SST is non-varying and zonally uniform. The longitudinal variation is specified using the “Qobs” SST pattern from
 827 Neale and Hoskins (2000), given by:

$$828 T(\varphi) = \begin{cases} \frac{1}{2} (2 - \sin^4 \phi - \sin^2 \phi) \delta T + T_{\min}, & \text{if } |\varphi| < \frac{\pi}{3} \\ 0, & \text{otherwise} \end{cases} \quad (\text{B1})$$

829 where φ is latitude, $\phi = \frac{\pi}{2} \frac{\varphi}{\varphi_{\max}}$, $\varphi_{\max} = \frac{\pi}{3}$, $\delta T = T_{\max} - T_{\min}$, $T_{\max} = 27^\circ\text{C}$, and $T_{\min} = 0^\circ\text{C}$.

830
 831 Because results are sensitive to the specification of the SSTs, groups that use a prognostic equation for the surface skin
 832 temperature are asked to set this skin temperature to the specified SST. No sea ice is prescribed, so the surface temperature is
 833 spatially uniform at 0°C poleward of 60° for the control simulation.

834 Radiatively active trace gases are well-mixed with mixing ratios following the AMIP II recommendations: CO₂: 348 ppmv;
 835 CH₄: 1650 ppbv; N₂O: 306 ppbv; Halocarbon yield of approximately 0.24 W m⁻² radiative forcing. The ozone distribution is
 836 the same as used in APE and CFMIP2/CMIP5, and is derived from the climatology used in AMIP II (Gates et al., 1999), and
 837 is constant in time and symmetric zonally and about the equator. This ozone distribution is provided as a netCDF file which is
 838 archived on the Earth System Grid and available via the DOI <http://dx.doi.org/10.5065/D61834Q6>. Ozone values are
 839 provided up to 0.28hPa (about 60km altitude in mid-latitudes). For models with tops above this level, a high top ozone
 840 dataset is also provided, which is available via the DOI <http://doi.org/10.5065/D64X5653>. The ozone climatologies provided
 841 uses pressure as a vertical coordinate. Most models use a sigma or hybrid vertical coordinate in pressure or altitude, which
 842 will mean that the pressure on a given model level varies in time, near the surface at the very least. Although the ozone
 843 climatology can be interpolated to the pressure of each model level as it varies in time within the model, for simplicity we
 844 recommend interpolating the ozone dataset onto the model vertical grid before the experiment is performed, and then
 845 specifying ozone values which are constant in time on each model level. This vertical interpolation will require a zonally
 846 symmetric climatology of pressure on model levels which is as consistent as possible with that expected in the aqua-control
 847 experiment. This could for example be produced by initially running a test version of the aqua-control experiment with an
 848 ozone climatology taken from a more realistic model configuration such as the AMIP DECK experiment.

849 Aerosols are removed to the extent possible to remove aerosol-radiation interaction (aka direct effects) and aerosol-cloud
850 interaction (aka indirect effects). No external surface emissions are to be prescribed. Models requiring aerosol for cloud
851 condensation should use a constant oceanic climatology that is symmetric about the equator and zonally. Alternatively,
852 models with the capability should set the cloud droplet and crystal numbers to $100 \cdot 10^6 \text{ m}^{-3}$ and $0.1 \cdot 10^6 \text{ m}^{-3}$, respectively (as
853 in Medeiros et al., 2016).

854 As in APE, it is recommended that the atmospheric dry mass be adjusted to yield a global mean of 101080 Pa. It is also
855 recommended to adopt the APE recommended values for geophysical constants, as listed in Table 2 of Williamson et al.
856 (2012).

857 The aqua-4K experiment follows the above protocol, but with SST derived by adding 4K to Eq. B1.

858 The aqua-4xCO₂ experiment replaces the CO₂ mixing ratio with 1392 ppmv. The SST is unchanged from the control
859 simulation (Eq. B1).

860 Model runs should be 10 years. We recommend discarding the initial spin up period of a few months.

863 **APPENDIX C: SST Pattern for CFMIP *amip-future4K/amipFuture* experiments**

864
865 The *amip-future4K* (formerly *amipFuture*) experiment is the same as the *amip* DECK experiment, except that the SSTs are
866 subject to a composite SST warming pattern derived from the CMIP3 coupled models. The patterned SST forcing dataset is
867 available in a netcdf file called `cfmip2_4k_patterned_sst_forcing.vn1.0.nc` which is available in the supplementary
868 information for this paper, and via the CFMIP website. This is a normalised multi-model ensemble mean of the ocean
869 surface temperature response pattern (the change in ocean surface temperature (TOS) between years 0-20 and 140-160, the
870 time of CO₂ quadrupling in the 1% runs) from thirteen CMIP3 AOGCMs (cccma, cnrm, gfdlcm20, gfdlcm21, gisser,
871 inmcm3, ipsl, miroc-medres, miub, mpi, mri, ncar-ccsm3, and ncar-pcm1.) Before computing the multi-model ensemble
872 mean, each model's TOS response was divided by its global mean and multiplied by 4. This guarantees that the pattern
873 information from all models is weighted equally and the global mean SST forcing is the same as in the uniform +4K
874 experiment. We have retained the SST forcing based on the CMIP3 coupled models because we consider it more important to
875 be able to compare CMIP5 and CMIP6 models forced with the same SST pattern than to use a pattern which is consistent
876 with, say, the CMIP5 coupled response.
877

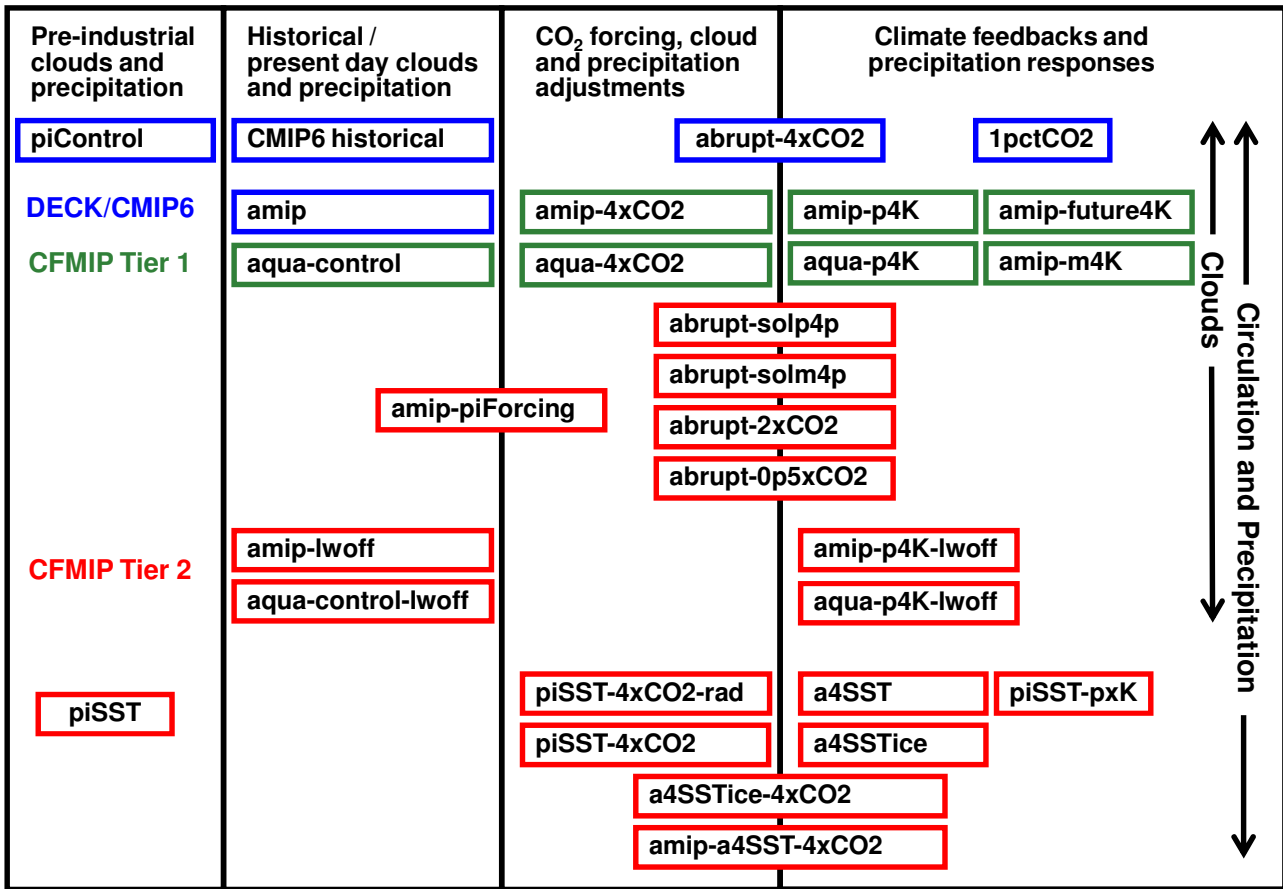
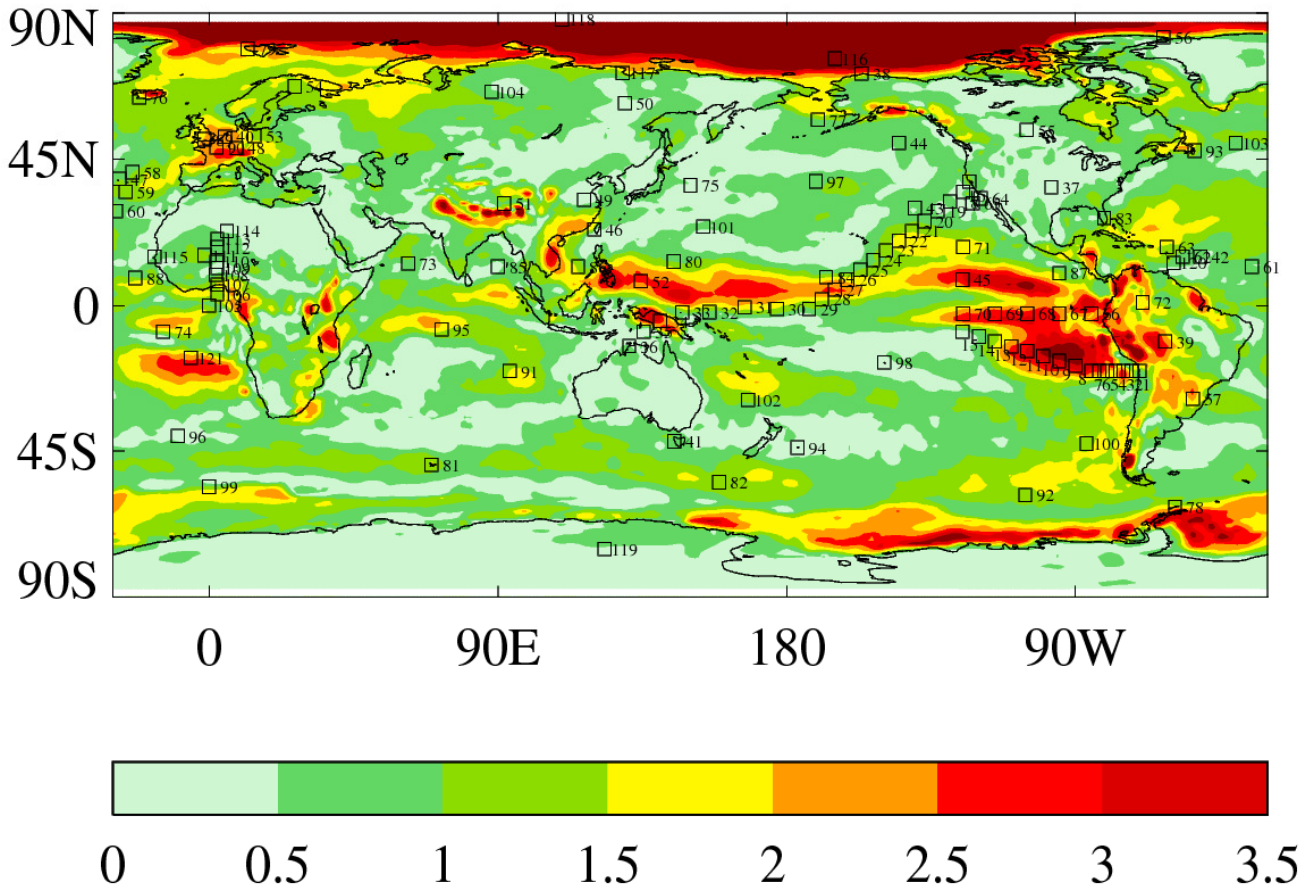


Figure 1. Summary of CFMIP-3 experiments and DECK + CMIP6 Historical experiments.

CFMIP-3 cfSites Locations



884
885
886
887

Figure 2. CFMIP-3 cfSites locations. The contours give an indication of inter-model spread in cloud feedback from the CFMIP-2 amip/amip4K experiments (please refer to Webb et al., 2015a for details).

888 **Table 1.** Summary of CFMIP Tier 1 experiments.

Experiment Name	Experiment Description / Design	Configuration	Start Year	Length
amip	This is a single ensemble member of the AMIP DECK experiment which contains additional outputs which are required for model evaluation using COSP, and as control values for model outputs in the amip-p4K, amip-4xCO2, amip-future4K and amip-m4K experiments.	Atmos-only	1979	36
amip-p4K	As CMIP5/CFMIP-2 amip4K experiment. AMIP experiment where SSTs are subject to a uniform warming of 4K.	Atmos-only	1979	36
amip-4xCO2	As CMIP5/CFMIP-2 amip4xCO2 experiment. AMIP experiment where SSTs are held at control values and the CO ₂ seen by the radiation scheme is quadrupled.	Atmos-only	1979	36
amip-future4K	As CMIP5/CFMIP-2 amipFuture experiment. AMIP experiment where SSTs are subject to a composite SST warming pattern derived from coupled models, scaled to an ice-free ocean mean of 4K.	Atmos-only	1979	36
aqua-control	Extended version of CMIP5/CFMIP-2 aquaControl experiment. Aquaplanet (no land) experiment with no seasonal cycle forced with specified zonally symmetric SSTs.	Atmos-only	1979	10
aqua-p4K	Extended version of CMIP5/CFMIP-2 aqua4K experiment. Aquaplanet experiment where SSTs are subject to a uniform warming of 4K.	Atmos-only	1979	10
aqua-4xCO2	Extended version of CMIP5/CFMIP-2 aqua4xCO2 experiment. Aquaplanet experiment where SSTs are held at control values and the CO ₂ seen by the radiation scheme is quadrupled.	Atmos-only	1979	10

889
890

891 **Table 2.** Summary of CFMIP Tier 2 experiments.

Experiment Name	Experiment Description / Design	Configuration	Start Year	Length
amip-m4K	As amip experiment but SSTs are subject to a uniform cooling of 4K.	Atmos-only	1979	36
amip-lwoff	As amip experiment, but with cloud-radiative effects switched off in the LW radiation code.	Atmos-only	1979	36
amip-p4K-lwoff	As amip-p4K experiment, but with cloud-radiative effects switched off in the LW radiation code.	Atmos-only	1979	36
aqua-control-lwoff	As aqua-control experiment, but with cloud-radiative effects switched off in the LW radiation code.	Atmos-only	1979	10
aqua-p4K-lwoff	As aqua-p4K experiment, but with cloud-radiative effects switched off in the LW radiation code.	Atmos-only	1979	10
abrupt-solp4p	Conceptually similar to abrupt 4xCO ₂ DECK experiment, except that the solar constant rather than CO ₂ is abruptly increased by 4%.	Coupled AOGCM	1850	150
abrupt-solm4p	Same as abrupt-solp4p, except solar constant is reduced by 4% rather than increased.	Coupled AOGCM	1850	150
abrupt-2xCO ₂	Identical to the DECK abrupt4xCO ₂ , but at 2xCO ₂ .	Coupled AOGCM	1850	150
abrupt-0p5xCO ₂	Identical to the DECK abrupt4xCO ₂ , but at 0.5xCO ₂	Coupled AOGCM	1850	150
amip-piForcing	Identical to AMIP DECK experiment but from 1870-present with constant pre-industrial forcing levels (anthro & natural).	Atmos-only	1870	145
piSST	An AGCM experiment with monthly-varying SSTs, sea-ice, atmospheric constituents and any other necessary boundary conditions (e.g. vegetation if required) taken from each model's own piControl run (using the 30 years of piControl that are parallel to years 111-140 of its abrupt4xCO ₂ run). Dynamic vegetation should be turned off in all the piSST set of experiments.	Atmos-only	Year 111 of abrupt-4xCO ₂	30
piSST-pxK	Same as piSST, but with a spatially and temporally uniform SST anomaly applied on top of the monthly-varying piSST SSTs. The magnitude of the uniform increase is taken from each model's global, climatological annual mean open SST change between abrupt4xCO ₂ minus piControl (using the mean of years 111-140 of abrupt4xCO ₂ , and the parallel 30-year section of piControl).	Atmos-only	Year 111 of abrupt-4xCO ₂	30
piSST-4xCO ₂ -rad	Same as piSST but CO ₂ as seen by the radiation scheme is quadrupled.	Atmos-only	Year 111 of abrupt-4xCO ₂	30
piSST-4xCO ₂	Same as piSST but CO ₂ is quadrupled. The increase in CO ₂ is seen by both the radiation scheme and vegetation.	Atmos-only	Year 111 of abrupt-4xCO ₂	30
a4SST	As piSST, but with monthly-varying SSTs taken from years 111-140 of each model's own abrupt4xCO ₂ experiment instead of from piControl. Sea-ice is unchanged from piSST.	Atmos-only	Year 111 of abrupt-4xCO ₂	30
a4SSTice	As piSST, but with monthly-varying SSTs and sea-ice taken from years 111-140 of each model's own abrupt4xCO ₂ experiment instead of from piControl.	Atmos-only	Year 111 of abrupt-4xCO ₂	30
a4SSTice-4xCO ₂	As a4SSTice, but CO ₂ is quadrupled, and the increase in CO ₂ is seen by both the radiation scheme and vegetation.	Atmos-only	Year 111 of abrupt-4xCO ₂	30
amip-a4SST-4xCO ₂	Same as amip, but a patterned SST anomaly is applied on top of the monthly-varying amip SSTs. This anomaly is a monthly climatology, taken from each model's own abrupt4xCO ₂ run minus piControl (using the mean of years 111-140 of abrupt4xCO ₂ , and the parallel 30-year section of piControl). CO ₂ is quadrupled, and the increase in CO ₂ is seen by both the radiation scheme and vegetation.	Atmos-only	1979	36

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Dataset	Years	Observables	Applications	References
CALIPSO-GOCCP	2006/06 - 2012/10	Cloud fractions: 2D and 3D by phase. Scattering ratio histograms as function of height.	Vertical distributions of clouds. Cloud phase identification.	Chepfer et al., (2010); Cesana and Chepfer, (2013)
CloudSat	2006/06 - 2010/12	Reflectivity histograms as function of height.	Vertical distributions of clouds and precipitation	Marchand et al., (2009); Zhang et al., (2010)
ISCCP	1983/07-2008/06	Cloud top pressure – cloud optical depth histograms.	Cloud radiative properties. Long time series.	Rossow and Schiffer, (1999)
MODIS	2002/07 – 2015/11	Cloud top pressure – cloud optical depth histograms. Total, liquid and ice cloud fractions. Effective radius – optical depth histograms by cloud phase.	Cloud radiative properties. Effective size, and phase information.	Pincus et al., (2012); King et al., (2003)
MISR	2000/06 – 2013/05	Cloud top height (CTH) – cloud optical depth histograms	Cloud radiative properties. Independent estimate of cloud top height.	Marchand et al., (2010)
PARASOL	2003/05 - 2012/08	Monodirectional reflectance	Cloud radiative properties.	Konsta et al., (2015)

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