

**Response to Topical Editor Decision: Publish subject to minor revisions (Editor review)**  
**by J. C. Hargreaves on “The Cloud Feedback Model Intercomparison Project (CFMIP) contribution to CMIP6” by Mark J. Webb et al.**

Editor comments below are shown in bold and our responses are in italics.

*Dear Julia,*

**Thanks for the revised manuscript and thorough responses to all the reviews and comments. I just have a few mostly technical questions.**

*Thank you.*

**The nomenclature in regards to the whole project is a bit inconsistent. It is stated that “CFMIP-3” is, in this paper, to apply to the CFMIP-3 in CMIP6 runs, while also acknowledging that there will be other CFMIP-3 runs that are not included in CMIP6. What are they to be called? At the same time, throughout the paper reference is made to “CMIP5/CFMIP-2”, which presumably refers to the subset of CFMIP-2 runs that were included in CMIP5. Therefore I think it might be more consistent to use “CMIP6/CFMIP-3” in this manuscript, particularly in places where a comparison is being made with “CMIP5/CMIP-2”.**

*We agree that this is inconsistent. We think the best way to address this is to refer to the broader CFMIP-3 project as CFMIP-3, but to refer to the various experiments as ‘CFMIP-2/CMIP5 experiments’, ‘CFMIP-3/CMIP6 experiments’ and ‘informal CFMIP-3 experiments’. We have amended the manuscript throughout. We have also rewritten the relevant text in the introduction thus:*

*“CFMIP is now entering its third phase, CFMIP-3, which will run in parallel with the current phase of the Coupled Model Intercomparison Project (CMIP6, Eyring et al., 2016) This paper documents the CFMIP-3/CMIP6 experiments and diagnostic outputs which constitute the CFMIP-3 contribution to CMIP6. It is anticipated that CFMIP-3 will be broader than what is described here, for instance including studies with process models, and informal CFMIP-3 experiments which are organised independently of CMIP6. Please refer to the CFMIP website for announcements of these other initiatives and CFMIP annual meetings.”*

**The nonlinearity of the Gregory plot is alluded to by Anonymous reviewer 1, and you responded, but do not include this discussion in the manuscript. It seems to me that it should be added, as otherwise the paper may give many people the impression that the estimation of climate sensitivity by the Gregory method is wholly accurate.**

*We have added the following to the end of Section 2.7:*

*“We also consider the time variation of feedbacks in abrupt-4xCO<sub>2</sub> experiments to be an important area to be investigated, as this can have a substantial impact on estimates of equilibrium sensitivity*

*(e.g. Geoffroy et al., 2013). Andrews et al., 2015 investigated such effects using two atmosphere-only GCMs forced with SSTs and sea ice from their own abrupt-4xCO2 experiments, and attributed the time variation in the feedbacks to changes in the pattern of surface warming. Pilot studies are ongoing to develop similar experiments based on a composite SST pattern response more representative of the CMIP5 ensemble mean. We plan to organise an informal pilot intercomparison based on this within CFMIP-3 and may subsequently propose these experiments as an extension to the CFMIP-3/CMIP6 experiment set.”*

**It seems a bit inconsistent that discussion of Aiko’s suggestion for an experiment has been added to the paper, but the suggestions made by Anonymous reviewers 1 and 2 have not been mentioned. What is the reason for this?**

*Anonymous reviewers 1 and 2 both proposed new experiments, but Aiko proposed doing existing experiments slightly differently. We felt that it was appropriate to mention Aiko’s points in the text because they are relevant to the design of the experiments which we *\*are\** proposing for CFMIP-3/CMIP6, as opposed to the suggestions from anonymous reviewers which are about experiments which we *\*aren’t\** proposing.*

**It is a shame that it is not anymore possible to change the names of the individual runs as, if one reviewer is confused by the naming, then other people will be too.**

*Yes, we agree this is unfortunate. However the description of the experiment is clear. To minimise potential confusion we have added the following to the manuscript:*

*“(Given the names of other CMIP6 experiments this might have been better named amip-4xCO2-rad, but this inconsistency was only noticed after the experiment names were finalised and propagated to the ESGF).”*

*Regards,*

*Mark Webb*

*(On behalf of the authors)*

# The Cloud Feedback Model Intercomparison Project (CFMIP) contribution to CMIP6.

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## Abstract

The primary objective of CFMIP is to inform future assessments of cloud feedbacks through improved understanding of cloud-climate feedback mechanisms and better evaluation of cloud processes and cloud feedbacks in climate models. However, the CFMIP approach is also increasingly being used to understand other aspects of climate change, and so a second objective has now been introduced, to improve understanding of circulation, regional-scale precipitation, and non-linear changes. CFMIP is supporting ongoing model inter-comparison activities by coordinating a hierarchy of targeted experiments for CMIP6, along with a set of cloud related output diagnostics. CFMIP contributes primarily to addressing the CMIP6 questions "How does the Earth System respond to forcing?" and "What are the origins and consequences of systematic model biases?" and supports the activities of the WCRP Grand Challenge on Clouds, Circulation and Climate Sensitivity.

A compact set of Tier 1 experiments is proposed for CMIP6 to address the question: "1) What are the physical mechanisms underlying the range of cloud feedbacks and cloud adjustments predicted by climate models, and which models have the most credible cloud feedbacks?" Additional Tier 2 experiments are proposed to address the following questions: 2) Are cloud feedbacks consistent for climate cooling and warming, and if not, why? 3) How do cloud-radiative effects impact the structure, the strength and the variability of the general atmospheric circulation in present and future climates? 4) How do responses in the climate system due to changes in solar forcing differ from changes due to CO<sub>2</sub>, and is the response sensitive to the sign of the forcing? 5) To what extent is regional climate change per CO<sub>2</sub> doubling state-dependent (nonlinear), and why? 6) Are climate feedbacks during the 20<sup>th</sup> century different to those acting on long term climate change and climate sensitivity? 7) How do regional climate responses (e.g. in precipitation) and their uncertainties in coupled models arise from the combination of different aspects of CO<sub>2</sub> forcing and sea surface warming?

CFMIP also proposes a number of additional model outputs in the CMIP DECK, CMIP6 Historical and CMIP6 CFMIP experiments, including COSP simulator outputs and process diagnostics to address the following questions: 1) How well do clouds and other relevant variables simulated by models agree with observations? 2) What physical processes and mechanisms are important for a credible simulation of clouds, cloud feedbacks and cloud adjustments in climate models? 3) Which models have the most credible representations of processes relevant to the simulation of clouds? 4) How do clouds and their changes interact with other elements of the climate system?

## 1 Introduction

Inter-model differences in cloud feedbacks continue to be the largest source of uncertainty in predictions of equilibrium climate sensitivity (Boucher et al., 2013). Although the ranges of cloud feedbacks and climate sensitivity from comprehensive climate models have not reduced in recent years, considerable progress has been made in understanding (a) which types of clouds contribute most to this spread (e.g. Bony and Dufresne 2005; Webb et al., 2006; Zelinka et al., 2013), (b) the role of cloud adjustments in climate sensitivity (e.g. Gregory and Webb, 2008; Andrews and Forster, 2008; Kamae and Watanabe, 2012; Zelinka et al., 2013), (c) the processes and mechanisms which are (and are not) implicated in cloud

57 feedbacks, both in fine resolution models (e.g. Rieck et al., 2012; Bretherton et al., 2015) and in comprehensive climate  
58 models (e.g. Brient and Bony 2012; Sherwood et al., 2014; Zhao, 2015; Webb et al., 2015b), (d) the inconstancy of cloud  
59 feedbacks and effective climate sensitivity (e.g. Senior and Mitchell, 2000; Williams et al., 2008; Andrews et al., 2012;  
60 Geoffroy et al., 2013; Armour et al., 2013; Andrews and Gregory, 2016) and (e) the extent to which models with stronger or  
61 weaker cloud feedbacks or climate sensitivities agree with observations (e.g. Fasullo and Trenberth, 2012; Su et al., 2014; Qu  
62 et al., 2014; Sherwood et al., 2014; Myers and Norris, 2016). Additionally, our ability to evaluate model clouds using  
63 satellite data has benefited from the increasing use of satellite simulators. This approach, first introduced by Yu et al., 1996  
64 for use with data from the International Satellite Cloud Climatology Project (ISCCP) attempts to reproduce what a satellite  
65 would observe given the model state. Such approaches enable more quantitative comparisons to the satellite record (e.g. Yu  
66 et al., 1996; Klein and Jakob, 1999; Webb et al., 2001; Bodas-Salcedo et al., 2011; Cesana and Chepfer, 2013). Much of our  
67 improved understanding in these areas would have been impossible without the continuing investment of the scientific  
68 community in successive phases of the Coupled Model Intercomparison Project (CMIP), and its co-evolution in more recent  
69 years with the Cloud Feedback Model Intercomparison Project (CFMIP).

70 CFMIP started in 2003 and its first phase (CFMIP-1) organised an intercomparison based on perpetual July SST forced  
71 Cess style +2K experiments and 2xCO<sub>2</sub> equilibrium mixed-layer model experiments containing ISCCP simulator in parallel  
72 with CMIP3 (McAvaney and Le Treut, 2003). CFMIP-1 had a substantial impact on the evaluation of clouds in models and  
73 in the identification of low level cloud feedbacks as the primary cause of inter-model spread in cloud feedback, which  
74 featured prominently in the fourth and fifth IPCC assessments (Randall et al., 2007; Boucher et al., 2013).

75 The subsequent objective of CFMIP-2 was to inform improved assessments of climate change cloud feedbacks by  
76 providing better tools to support evaluation of clouds simulated by climate models and understanding of cloud-climate  
77 feedback processes. CFMIP-2 organized further experiments as part of CMIP5 (Bony et al., 2011; Taylor et al., 2012),  
78 introducing seasonally varying SST perturbation experiments for the first time, as well as fixed SST CO<sub>2</sub> forcing experiments  
79 to examine cloud adjustments. CFMIP-2 also introduced idealized 'aquaplanet' experiments into the CMIP family of  
80 experiments. These experiments were motivated by extensive research in the framework of the aqua-planet experiment  
81 (Neale and Hoskins, 2000, Blackburn and Hoskins, 2013) and the particular finding, based on a small subset of models, that  
82 the global mean cloud feedback of more realistic model configurations could be reproduced, and more easily investigated,  
83 using the much simpler aqua-planet configuration (Medeiros et al., 2008). CFMIP-2 proposed the inclusion of the abrupt  
84 CO<sub>2</sub> quadrupling AOGCM (atmosphere-ocean general circulation model) experiment in the core experiment set of CMIP5,  
85 based on the approach of Gregory et al., 2004, which subsequently formed the basis for equilibrium climate sensitivity  
86 estimates from AOGCMs (Andrews et al., 2012). Additionally CFMIP-2 introduced satellite simulators to CMIP via the  
87 CFMIP Observation Simulator Package (COSP, Bodas-Salcedo et al., 2011); not only the ISCCP simulator, but additional  
88 simulators to facilitate the quantitative evaluation clouds using a new generation of active radars and lidars in space. CFMIP-  
89 2 also introduced into CMIP5 process diagnostics such as temperature and humidity budget tendency terms and high  
90 frequency 'cfSites' outputs at 120 locations around the globe. In an effort less directly connected to CMIP, CFMIP organized  
91 a joint project with the GEWEX Global Atmospheric System Study (GASS) called CGILS (the CFMIP-GASS  
92 Intercomparison of LES and SCMs) to develop cloud feedback intercomparison cases to assess the physical credibility of  
93 cloud feedbacks in climate models by comparing Single Column Model (SCM) versions of General Circulation Models  
94 (GCMs) with high resolution Large Eddy Simulations (LES) models. CFMIP-2 also developed the CFMIP-OBS data portal  
95 and the CFMIP diagnostic codes catalogue. For more details, and for a full list of CFMIP related publications, please refer  
96 to the CFMIP website (<http://www.earthsystemcog.org/projects/cfmip>).

97 Studies arising from CFMIP-2 include numerous single and multi-model evaluation studies which use COSP to make  
98 quantitative and fair comparisons with a range of satellite products (e.g. Kay et al., 2012; Franklin et al., 2013; Klein et al.,  
99 2013, Lin et al., 2014, Chepfer et al., 2014.). COSP has also enabled studies attributing cloud feedbacks and cloud  
100 adjustments to different cloud types (e.g. Zelinka et al., 2013; Zelinka et al., 2014; Tsushima et al., 2015). CFMIP-2  
101 additionally enabled the finding that idealized 'aquaplanet' experiments without land, seasonal cycles or Walker circulations  
102 are able to reproduce the essential differences between models' global cloud feedbacks and cloud adjustments in a substantial  
103 ensemble of models (Ringer et al., 2014; Medeiros et al., 2015). Process outputs from CFMIP have also been used to develop  
104 and test physical mechanisms proposed to explain and constrain inter-model spread in cloud feedbacks in the CMIP5 models  
105 (e.g. Sherwood et al., 2014; Brient et al., 2015; Webb et al., 2015a; Nuijens et al., 2015a,b; Dal Gesso et al., 2015). CGILS  
106 has demonstrated a consensus in the responses of LES models to climate forcings and identified shortcomings in the physical  
107 representations of cloud feedbacks in climate models (e.g. Blossey et al., 2013; Zhang et al., 2013; Dal Gesso et al., 2015).

108 The CFMIP experiments have additionally formed the basis for coordinated experiments to explore the impact of cloud  
109 radiative effects on the circulation (Stevens et al., 2012; Fermepin and Bony 2014; Crueger and Stevens 2015; Li et al., 2015;  
110 Harrop and Hartmann 2016), the impact of parametrized convection on cloud feedback (Webb et al., 2015b) and the  
111 mechanisms of negative shortwave cloud feedback in mid to high latitudes (Ceppi et al., 2015). Additionally the CFMIP  
112 experiments have, due to their idealized nature, proven useful in a number of studies not directly related to clouds, but instead  
113 analyzing the responses of regional precipitation and circulation patterns to CO<sub>2</sub> forcing and climate change (e.g. Bony et al.,  
114 2013; Chadwick et al., 2014; He and Soden 2015; Oueslati et al., 2016). Studies using CFMIP-2 outputs from CMIP5 remain  
115 ongoing and further results are expected to feed into future assessments of the representation of clouds and cloud feedbacks in  
116 climate models.

117 The primary objective of CFMIP is to inform future assessments of cloud feedbacks through improved understanding of  
118 cloud-climate feedback mechanisms and better evaluation of cloud processes and cloud feedbacks in climate models.  
119 However, the CFMIP approach is also increasingly being used to understand other aspects of climate change, and so a second  
120 objective has been introduced, to improve understanding of circulation, regional-scale precipitation, and non-linear changes.

121 This involves bringing climate modelling, observational and process modelling communities closer together and providing  
122 better tools and community support for evaluation of clouds and cloud feedbacks simulated by climate models and for  
123 understanding of the mechanisms underlying them. This is achieved by:

- 125 • Coordinating model inter-comparison activities which include experimental design as well as specification of  
126 model output diagnostics to support quantitative evaluation of modelled clouds with observations (e.g. COSP)  
127 and in-situ measurements (e.g. cfSites) as well as process-based investigation of cloud maintenance and  
128 feedback mechanisms (e.g. cfSites, temperature and humidity tendency terms)
- 129 • Developing and improving support infrastructure including COSP, CFMIP-OBS and the CFMIP diagnostic  
130 codes catalogue.
- 131 • Fostering collaboration with the observational and cloud process modelling communities via annual CFMIP  
132 meetings and international funded projects.

133 CFMIP is now entering its third phase, CFMIP-3, which will run in parallel with the current phase of the Coupled Model  
134 Intercomparison Project (CMIP6, Eyring et al., 2016). This paper documents the CFMIP-3/CMIP6 experiments and  
135 diagnostic outputs which constitute the CFMIP-3 contribution to CMIP6. It is anticipated that CFMIP-3 will be broader than  
136 what is described here, for instance including studies with process models, and informal CFMIP-3 experiments which are  
137 organised independently of CMIP6. Please refer to the CFMIP website for announcements of these other initiatives and  
138 CFMIP annual meetings.

139 CFMIP-3 touches, to differing degrees, on each of the three questions around which CMIP6 is organized. With its focus on  
140 cloud feedback, CFMIP-3 is central to CMIP6's attempt to answer the question: 'How does the Earth system respond to  
141 forcing?' But as illustrated in the remainder of this document, CFMIP-3 also offers the opportunity to contribute to the other  
142 two guiding questions of CMIP6. Through its strong model evaluation component it stands to help answer the question:  
143 'What are the origins and consequences of systematic model biases?' CFMIP-3 will also help answer the question: 'How can  
144 we assess future climate changes given climate variability, climate predictability, and uncertainties in scenarios?' For  
145 example the *amip-piForcing* experiment proposed below will support studies relating cloud variability and feedbacks on  
146 observable timescales to long term cloud feedbacks (Andrews, 2014; Gregory and Andrews, 2016).

147 The CFMIP-3/CMIP6 experiments are outlined below in Section 2. Section 3 describes the diagnostics outputs proposed  
148 by CFMIP for the CFMIP-3/CMIP6 experiments and other experiments within CMIP. We provide a summary of the  
149 CFMIP-3 contribution to CMIP6 in Section 5.

## 151 2 CFMIP-3 Experiments

152 The CFMIP-3/CMIP6 experiments are summarised in Figure 1 and Tables 1 and 2, and are described in detail below. Most of  
153 the CFMIP-3/CMIP6 experiments are based on CO<sub>2</sub> concentration forced amip, piControl and abrupt-4xCO<sub>2</sub> CMIP DECK  
154 (Diagnostic, Evaluation and Characterization of Klima) experiments (Eyring et al., 2016). Unless otherwise specified  
155 below, the CFMIP-3/CMIP6 experiments should be configured consistently with the DECK experiments on which they are  
156 based, using consistent model formulation, and forcings and boundary conditions as specified by Eyring et al., 2016.  
157 Following the CMIP6 design protocol, groups of experiments are motivated by science questions and are separated into Tiers  
158 1 and 2 (Eyring et al., 2016). It is a requirement for participation by modelling groups in the CFMIP-3/CMIP6 model  
159 intercomparison that all Tier 1 experiments be performed and published through the ESGF, so as to support CFMIP's Tier 1  
160 science question. Tier 2 experiments are optional, and are associated with additional science questions. Any subset of Tier  
161 2 experiments may be performed. All model output archived by CFMIP-3/CMIP6 is expected to be made available under the  
162 same terms as CMIP output. Most modelling groups currently release their CMIP data for unrestricted use. Our analysis  
163 plans for the CFMIP-3/CMIP6 experiments are summarised in Appendix A.

### 165 2.1 CFMIP-3/CMIP6 Tier 1 Experiments

166 Lead coordinator: Mark Webb

169 Science Question: What are the physical mechanisms underlying the range of cloud feedbacks and cloud adjustments  
170 predicted by climate models, and which of the cloud responses are the most credible?

172 Equilibrium climate sensitivity (ECS) can be estimated using an idealized AOGCM experiment such as the *abrupt-4xCO<sub>2</sub>*  
173 experiment in the CMIP6 DECK, at the same time statistically separating the global mean contributions from climate  
174 feedbacks and adjusted radiative forcing due to CO<sub>2</sub> (Gregory et al. 2004, Andrews et al., 2012). However understanding the  
175 physical processes underlying cloud feedbacks and adjustments requires diagnosis in SST forced experiments with  
176 atmosphere-only general circulation models (AGCMs), which can resolve cloud feedbacks and adjustments independently  
177 from each other and with minimal statistical noise at regional scales, while faithfully reproducing the inter-model differences  
178 in global values from the fully coupled models (Ringer et al., 2014). (The ability of these AGCM experiments to reproduce  
179 the inter-model differences in global cloud feedbacks and adjustments from coupled models indicates that they do not  
180 strongly depend on different ocean model formulations or SST biases). The CFMIP-2/CMIP5 *amip4xCO<sub>2</sub>* experiments  
181 which quadrupled CO<sub>2</sub> while leaving SSTs at present day values (Bony et al., 2011), allowed the land/tropospheric

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199 adjustment process and the cloud adjustment to CO<sub>2</sub> to be examined in this way for the first time in the multi-model context  
200 (Kamae and Watanabe, 2012; Ringer et al., 2014; Kamae et al. 2015) in conjunction with the CMIP5 *sstClim/sstClim4xCO2*  
201 experiments which were based on climatological preindustrial SSTs (Andrews et al., 2012; Zelinka et al., 2013; Vial et al.,  
202 2013). These experiments have additionally formed the basis for more in-depth studies with individual models (e.g. Wyant et  
203 al., 2012; Kamae and Watanabe, 2013; Bretherton et al., 2014, Ogura et al., 2014). The CFMIP-2/CMIP5 *amip4K* and  
204 *amipFuture* SST perturbed atmosphere-only experiments (Bony et al., 2011) have been used to examine cloud feedbacks in  
205 greater detail (e.g. Brient and Bony, 2012; Bretherton et al., 2014; Lacagnina et al., 2014; Bellomo and Clement, 2015; Webb  
206 et al., 2015b), often in conjunction with simulator outputs (e.g. Gordon and Klein, 2014; Chepfer et al., 2014; Tsushima et al.,  
207 2015, Ceppi et al., 2016) and CFMIP process diagnostics (e.g. Webb and Lock, 2013; Sherwood et al., 2014; Brient et al.,  
208 2015; Webb et al., 2015a; Dal Gesso et al., 2015). Similarly, these experiments have been used to investigate regional  
209 responses of various quantities to direct radiative forcing due to increasing CO<sub>2</sub> concentrations and/or increases in SST,  
210 including precipitation (e.g. Ma and Xie, 2013; Huang et al., 2013; Widlansky et al., 2013; Kent et al., 2015; Long et al.,  
211 2016), circulation (e.g. He et al., 2014; Zhou et al., 2014; Kamae et al., 2014; Bellomo and Clement, 2015; Shaw and Voigt,  
212 2015) and stability (e.g. Qu et al., 2015).

213 A more idealized set of fixed SST experiments proposed by CFMIP-2 for CMIP5 (*aquaControl*, *aqua4xCO2*, and  
214 *aqua4K*) based on zonally symmetric, fixed season ‘aquaplanet’ configurations without land have been shown to reproduce  
215 the inter-model differences in global mean cloud adjustments and feedbacks from realistic experiments surprisingly  
216 effectively (Medeiros et al., 2008; Ringer et al., 2014; Medeiros et al., 2015) as well as many aspects of the zonal mean  
217 circulation response (Medeiros et al., 2015). This indicates that those features of the climate system excluded from these  
218 experiments (i.e. the ocean, land, seasonal cycle, monsoon and Walker circulations) are not central to understanding inter-  
219 model differences in global mean cloud feedbacks and adjustments, and demonstrates the value of aquaplanet experiments for  
220 investigating the origin of such differences, as well as differences in zonally averaged precipitation and circulation and their  
221 responses to climate change (e.g. Stevens et al., 2012; Bony et al., 2013; Oueslati and Bellon, 2013; Fermepin and Bony  
222 2014; Voigt and Shaw 2015). The aquaplanet experiments have the benefit not only of being less computationally expensive  
223 than alternative experiments (requiring only 5-10 years to get a robust signal); they are also much more straightforward to  
224 analyse, as their behaviour can mostly be characterized by examining zonal means, avoiding the analysis overhead of  
225 compositing which is generally required in realistic model configurations to isolate the various cloud regimes. Aqua-planet  
226 simulations (and other idealized) experiments are particularly effective at highlighting model differences, for instance in the  
227 placement of the tropical rain bands, or in the representation of cloud changes with warming, as it is not possible to tune them  
228 to observations in the same way as is for more realistic configurations (e.g., Stevens and Bony, 2013).

229 The CFMIP-2/CMIP5 experiments and diagnostic outputs have thus enabled considerable progress on a number of  
230 questions. However, participation by a larger fraction of modelling groups is desired in CFMIP-3/CMIP6 to enable a more  
231 comprehensive assessment of the uncertainties across the full multi-model ensemble. Our proposal is therefore to retain the  
232 CFMIP-2/CMIP5 experiments (known in CMIP5 as *amip4K*, *amip4xCO2*, *amipFuture*, *aquaControl*, *aqua4xCO2* and  
233 *aqua4K*) in Tier 1 for CFMIP-3/CMIP6. These are summarised in Table 1 (the names have been changed slightly compared  
234 to the CMIP5 equivalents to fit in with a wider naming convention of CMIP6). The set up for each of these experiments is  
235 described below. (For output requirements from these and other experiments please refer to Section 3).

236 *amip*: This is a single ensemble member of the CMIP DECK *amip* experiment which contains additional outputs which are  
237 required both for model evaluation using COSP, and for interpretation of feedbacks and adjustments in conjunction with the  
238 *amip-p4K*, *amip-4xCO2*, *amip-future4K* and *amip-m4K* experiments.

239 *amip-p4K* (formerly *amip4K*): The same as the *amip* DECK experiment, except that SSTs are subject to a uniform  
240 warming of 4K. This warming should be applied to the ice free ocean surface only. Sea ice and SSTs in grid boxes  
241 containing sea ice remain the same as in the *amip* DECK experiment.

242 *amip-future4K* (formerly *amipFuture*): The same as the *amip* DECK experiment, except that a composite SST warming  
243 pattern derived from the CMIP3 coupled models is added to the AMIP SSTs (see Appendix C for details). As with the *amip*-  
244 *p4K* experiment, the warming pattern should only be applied to the ice free ocean surface, and sea ice and SSTs in grid boxes  
245 containing sea ice should remain the same as in the *amip* DECK experiment. The warming pattern should be scaled to ensure  
246 that the global mean SST increase averaged over the ice free oceans is 4K. Care should be taken to ensure that SSTs are  
247 increased in any inland bodies of water and near coastal edges, for example by linearly interpolating the provided warming  
248 pattern dataset to fill in missing data before re-gridding to the target resolution.

249 *amip-4xCO2* (formerly *amip4xCO2*): The same as the *amip* experiment within the DECK, except that the CO<sub>2</sub>  
250 concentration seen by the radiation scheme is quadrupled. The CO<sub>2</sub> seen by the vegetation should be the same as in the *amip*  
251 DECK experiment. This experiment gives an indication of the adjusted radiative forcing due to CO<sub>2</sub> quadrupling, including  
252 stratospheric, land surface, tropospheric and cloud adjustments. (Given the names of other CMIP6 experiments this  
253 experiment might have been better named *amip-4xCO2-rad*, but this inconsistency was only noticed after the experiment  
254 names were finalised and propagated to the ESGF).

255  
256 The configuration of the *aqua-control*, *aqua-p4K* and *aqua-4xCO2* experiments are unchanged compared to their  
257 equivalents in CFMIP-2/CMIP5, except that the simulation length has been extended to 10 years to improve the signal to  
258 noise ratio. Further details of their experimental set up are included in Appendix B.

259 We also propose to use the Tier 1 experiments as the foundation for further experiments planned in the context of the  
260 Grand Challenge on Clouds, Circulation and Climate Sensitivity (Bony et al., 2015). These will include for example  
261 sensitivity experiments to assess the impacts of different physical processes on cloud feedbacks and regional  
262 circulation/precipitation responses and also to test specifically proposed cloud feedback mechanisms (e.g. Webb et al., 2015b,

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264 Ceppi et al., 2015). Additional experiments further idealizing the aquaplanet framework to a non-rotating rotationally  
265 symmetric case are also under development (e.g. Popke et al., 2013). These will be proposed as additional Tier 2  
266 experiments at a future time, or coordinated by informally outside of CMIP6.

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## 267 **2.2 amip minus 4K Experiment (Tier 2)**

268 Lead Coordinators: Mark Webb and Bjorn Stevens

269 Science Question: Are cloud feedbacks consistent for climate cooling and warming, and if not, why?

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271  
272 There is some evidence to suggest that cloud feedbacks might operate differently in response to cooling rather than warming.  
273 For example, Yoshimori et al., 2009 found a positive shortwave cloud feedback in a CO<sub>2</sub> doubling experiment with a  
274 particular GCM, but noted a tendency for it to become weaker or even negative in cooling experiments designed to replicate  
275 the climate of the last glacial maximum. They suggested that this might be related to different displacements of mixed-phase  
276 clouds in the two scenarios. For small enough changes where linearity is a good approximation, one would expect the cloud  
277 response to cooling and warming to be the same, differing only in sign, resulting in an identical cloud feedback expressed per  
278 degree of global temperature change. But for larger perturbations this symmetry of response may no longer hold. A  
279 warming or cooling of the atmosphere of equal magnitude while maintaining relative humidity will for example generate  
280 different changes in absolute humidity, and its horizontal and vertical gradients, which have been linked to cloud feedbacks  
281 (Brient and Bony, 2013; Sherwood et al., 2014), the atmospheric lapse rate and circulation which influences clouds and  
282 depends in part on the absolute humidity (Held and Soden, 2006; Qu et al., 2015) and additionally on extratropical cloud  
283 optical depth feedbacks which may be related to adiabatic cloud liquid water contents (Gordon and Klein, 2014) or phase  
284 changes that depend upon whether a given volume crosses the 0 degree isotherm in the climate change (Ceppi et al. 2015).

285 The configuration of the *amip-m4K* experiment will be the same as the *amip-p4K* experiment, except that the sea surface  
286 temperatures are uniformly reduced by 4K rather than increased. This cooling should be applied to sea ice free grid boxes  
287 only. Sea ice and SSTs in grid boxes containing sea ice should remain the same as in the *amip* DECK experiment. In models  
288 which employ a fixed lower threshold near freezing for the SST used in the calculation of the surface fluxes, this should  
289 ideally also be reduced by 4K. This experiment will contain CFMIP COSP and process outputs so as to support the  
290 investigation of inconsistent responses of clouds to a cooling vs. a warming climate in a controlled way through comparison  
291 with the *amip-p4K* experiment. This experiment also complements the abrupt 0.5xCO<sub>2</sub> and the -4% solar experiments in that  
292 one can identify asymmetries in the warming/cooling response with and without interactions with the ocean. As such we hope  
293 that these experiments will provide useful synergies with the Palaeoclimate Model Intercomparison Project (PMIP) CMIP6  
294 experiments (Kageyama et al., 2016), for example in interpreting differing cloud feedbacks between future CO<sub>2</sub> forced  
295 experiments and those representing the Last Glacial Maximum, as highlighted by Yoshimori et al., 2009.

## 297 **2.3 Atmosphere-only experiments without longwave cloud radiative effects. (Tier 2)**

298 Lead Coordinators: Sandrine Bony and Bjorn Stevens

299 Science question: How do cloud-radiative effects impact the structure, the strength and the variability of the general  
300 atmospheric circulation in present and future climates?

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302  
303 It is increasingly recognized that clouds, and atmospheric cloud-radiative effects in particular, play a critical role in the  
304 general circulation of the atmosphere and its response to global warming or other perturbations: they have been found to  
305 modulate the structure, the position and shifts of the ITCZ (e.g. Slingo and Slingo 1988; Randall et al., 1989; Sherwood et al  
306 1994; Bergman and Hendon 2000; Hwang and Frierson, 2013; Fermepin and Bony 2014; Voigt et al., 2014; Loeb et al.,  
307 2015), the organisation of convection in tropical waves, Madden-Julian Oscillations and other forms of convective  
308 aggregation (e.g. Lee et al., 2001; Lin and Mapes, 2004; Bony and Emanuel, 2005; Zurovac-Jevtic et al., 2006; Crueger and  
309 Stevens, 2015; Muller and Bony, 2015), the extra-tropical circulation and the position of eddy-driven jets (e.g. Ceppi et al.,  
310 2012; Ceppi et al., 2014; Grise and Polvani 2014; Li et al., 2015; Voigt and Shaw, 2015), and modes of interannual to  
311 decadal climate variability (e.g. Bellomo et al., 2015; Rädel et al., 2016; Yuan et al., 2016). A better assessment of this role  
312 would greatly help to interpret model biases (how much do biases in cloud-radiative properties contribute to biases in the  
313 structure of the ITCZ, in the position and strength of the storm tracks, in the lack of intra-seasonal variability, etc) and to  
314 inter-model differences in simulations of the current climate and in climate change projections (especially changes in regional  
315 precipitation and extreme events). More generally, a better understanding of how clouds couple to the circulation is expected  
316 to improve our ability to answer the four science questions raised by the WCRP Grand Challenge on Clouds, Circulation and  
317 Climate Sensitivity (Bony et al., 2015).

318 These questions provided the scientific motivation for the Clouds On/Off Klima Intercomparison Experiment (COOKIE)  
319 project proposed by the European consortium EUCLIPSE and CFMIP (Stevens et al., 2012). The COOKIE experiments,  
320 which have been run by four to eight climate models (depending on the experiment), switched off the cloud-radiative effects  
321 (clouds seen by the radiation code -and the radiation code only- were artificially made transparent) in an atmospheric model  
322 forced by prescribed SSTs. By doing so, the atmospheric circulation could feel the lack of cloud-radiative heating within the  
323 atmosphere, but the land surface could also feel the lack of cloud shading, which led to changes in land surface temperatures

325 and land-sea contrasts. The change in circulation between On and Off experiments resulted from both effects, obscuring to  
326 some degree the mechanisms through which the atmospheric cloud-radiative effects interact with the circulation for given  
327 surface boundary conditions. As the longwave cloud-radiative effects are felt mostly within the troposphere (representing  
328 most of the net atmospheric cloud-radiative heating) while the shortwave effects are felt mostly at the surface (e.g. L'Ecuyer  
329 and McGarragh 2010; Haynes et al., 2013), we could better isolate the role of tropospheric cloud-radiative effects on the  
330 circulation by running atmosphere-only experiments in which clouds are made transparent to radiation only in the longwave.  
331 In this configuration, the models will have a shortwave cloud feedback but no longwave cloud feedback. We note that the  
332 presence of clouds does affect the shortwave radiative heating of the atmosphere, although this is a much smaller effect than  
333 its longwave equivalent (e.g. Pendergrass and Hartmann, 2014).

334 Therefore we propose in Tier 2 a set of simple experiments similar to the *amip*, *amip-p4K*, *aqua-control* and *aqua-p4K*  
335 experiments within Tier 1, but in which cloud-radiative effects are switched off in the longwave part of the radiation code  
336 while retaining those in the shortwave (Fermepin and Bony, 2014). Care should also be taken to remove the effects of cloud  
337 on any longwave cooling used in other model schemes (e.g. turbulent mixing) if these are calculated independently of the  
338 radiation scheme. These experiments will be referred to as *amip-lwoff*, *amip-p4K-lwoff*, *aqua-control-lwoff* and *aqua-p4K-*  
339 *lwoff*. The analysis of idealized (aqua-planet) experiments will allow us to assess the robustness of the impacts found in more  
340 realistic (AMIP) configurations. It will also facilitate the interpretation of the results using simple dynamical models or  
341 theories, in collaboration with large-scale dynamicists (e.g. DynVar). The comparison of the inter-model spread of  
342 simulations between the standard and 'lwoff' experiments for present-day and warmer climates will help to identify which  
343 aspects of the inter-model spread depend on the representation of cloud-radiative effects, and which aspects do not, thus  
344 better highlighting other sources of spread. An alternative method (proposed by Aiko Voigt) was also considered, in which  
345 clear-sky heating rates would be applied in the atmosphere while retaining the all-sky fluxes at the surface. Although this  
346 approach would potentially isolate the effects of cloud heating in the atmosphere more cleanly than the lwoff experiments  
347 proposed here, it is yet to be demonstrated in a pilot study, and is considered more technically difficult to implement than the  
348 lwoff experiments, which are very similar to those piloted by Fermepin and Bony, 2014.

## 349 **2.4 Abrupt +/-4% Solar Forced AOGCM experiments (Tier 2)**

350 Lead coordinators: Chris Bretherton, Roger Marchand, Bjorn Stevens

351  
352 Science Question: How do responses in the climate system due to changes in solar forcing differ from changes due to CO<sub>2</sub>,  
353 and is the response sensitive to the sign of the solar forcing?  
354

355 While rapid adjustments in clouds and precipitation can easily be separated from conventional feedbacks in SST forced  
356 experiments, such a separation in coupled models is complicated by various issues, including the response of the ocean on  
357 decadal timescales. A number of studies have examined cloud feedbacks in coupled models subject to a solar forcing, which  
358 is generally associated with much smaller global cloud and precipitation adjustment, due to a smaller atmospheric absorption  
359 for a given top of atmosphere forcing (e.g. Lambert and Faull, 2007; Andrews et al., 2010), but the regional cloud and  
360 precipitation changes have yet to be rigorously investigated across models. Solar forcing also differs from greenhouse  
361 forcing through its different fingerprint on the vertical structure of warming (Santer et al., 2013) and small changes in the  
362 radiative heating near the tropopause may project measurably on tropospheric climate (e.g., Butler et al., 2010), for instance  
363 by influencing the baroclinicity in the upper troposphere and thus the storm-tracks (Bony et al., 2015).

364 A +4% solar experiment *abrupt-solp4p* is proposed which is analogous to the *abrupt-4xCO2* experiment but rather than  
365 changing CO<sub>2</sub> it would abruptly increase the solar constant by four percent and keep it fixed for 150 years, resulting in a  
366 global mean radiative forcing of a similar magnitude to that due to CO<sub>2</sub> quadrupling. When changing the solar constant, the  
367 shape of the spectral solar irradiance distribution should remain consistent with that in the piControl experiment. This  
368 experiment complements the DECK *abrupt-4xCO2* experiment, tests the forcing feedback framework for analyzing climate  
369 change, and would support our understanding of regional responses of the coupled system with and without CO<sub>2</sub> adjustments.  
370 The complementary -4% abrupt solar forcing experiment (*abrupt-solm4p*) would allow the examination of feedback  
371 asymmetry under climate cooling, and would also help with the interpretation of model responses to geo-engineering  
372 scenarios and volcanic forcing, and of past climate signals.

## 373 **2.5 nonLinMIP abrupt 2xCO<sub>2</sub> and abrupt 0.5xCO<sub>2</sub> Experiments (Tier 2)**

374 Lead Coordinator: Peter Good

375  
376 Science Question: To what extent is regional-scale climate change per CO<sub>2</sub> doubling state-dependent (nonlinear); what are  
377 the associated mechanisms; and how does this affect our understanding of climate model uncertainty?  
378

379 Recent studies with individual, or a small number of climate models, have found substantial nonlinearities in regional-scale  
380 precipitation change (Good et al., 2012; Chadwick and Good, 2013), associated with robust physical mechanisms (Chadwick  
381 and Good, 2013). Significant nonlinearity has also been found in global and regional-scale warming (e.g. Colman and  
382 McAvaney, 2009; Jonko et al., 2013; Good et al., 2015; Meraner et al., 2013) and ocean heat uptake (Bouttes et al., 2015).



383 To address this science question we propose two new experiments for Tier 2, *abrupt 2xCO<sub>2</sub>* and *abrupt 0.5xCO<sub>2</sub>*. These are  
384 the same as the DECK *abrupt4xCO<sub>2</sub>* experiment except that CO<sub>2</sub> concentrations are doubled and halved respectively relative  
385 to the preindustrial control. These experiments are based on a proven analysis approach, including traceability of these  
386 experiments to transient-forcing simulations (Good et al., 2016), to explore global and regional-scale nonlinear responses,  
387 highlighting different behaviour under business-as-usual scenarios, mitigation scenarios and palaeoclimate simulations.  
388 Additionally comparisons of the abrupt 2xCO<sub>2</sub> and abrupt 4xCO<sub>2</sub> experiments will help to establish the extent to which the  
389 latter accurately estimates the equilibrium climate sensitivity to CO<sub>2</sub> doubling (e.g. Gregory et al., 2004, Block and  
390 Mauritsen, 2013). Additional experiments (Good et al., 2016) may be proposed for Tier 2 in the future, or coordinated  
391 informally by CFMIP-3 outside of CMIP6. These include 100-year extensions to *abrupt-4xCO<sub>2</sub>* and *abrupt-2xCO<sub>2</sub>*; a 1%  
392 ramp-down from the end of the *1pctCO<sub>2</sub>* experiment; an abrupt step-down to 1xCO<sub>2</sub> from year 100 of the *abrupt-4xCO<sub>2</sub>*.  
393 These would be used to explore longer-timescale responses, quantify nonlinear mechanisms more precisely and understand  
394 the reversibility of climate change.

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## 395 2.6 Feedbacks in AMIP experiments (Tier 2)

396 Lead Coordinator: Timothy Andrews

397  
398 Science question: Are climate feedbacks during the 20<sup>th</sup> century different to those acting on long term climate change?  
399

400 Recent studies have shown significant time variation in climate feedbacks in response to CO<sub>2</sub> quadrupling (e.g. Andrews et  
401 al., 2012; Geoffroy et al., 2013; Armour et al., 2013; Andrews et al., 2015). This raises the possibility that feedbacks during  
402 the 20<sup>th</sup> century may be different to those acting on long term change, and hence has the potential to alleviate the apparent  
403 discrepancy between estimates of climate sensitivity from comprehensive climate models and from simple climate models  
404 fitted to observed warming trends (Collins et al., 2013). For example Gregory and Andrews, 2016 found that two models  
405 forced with observed monthly 20<sup>th</sup> century SST and sea-ice variations simulated effective climate sensitivities of about 2K,  
406 whereas these same models forced with patterns of long term SST change simulated effective climate sensitivities of over 3K  
407 and 4K.

408 The previous CFMIP-2/CMIP5 design was unable to diagnose the time-variation of feedbacks of explicit relevance to the  
409 historical period, because this requires the removal of the time varying forcing. To address this we propose an additional  
410 experiment called '*amip-piForcing*' (*amip* pre-industrial forcing) following the design of Andrews 2014 and Gregory and  
411 Andrews, 2016. This experiment is the same as the *amip* DECK experiment (i.e. using observed monthly updating SSTs and  
412 sea-ice), but run for the period 1870-present and with constant pre-industrial forcings (i.e. all anthropogenic and natural  
413 forcing boundary conditions identical to the *piControl* experiment). Since the forcing constituents do not change in this  
414 experiment it readily allows a simple diagnosis of the simulated atmospheric feedbacks to observed SST and sea-ice changes,  
415 which can then be compared to feedbacks representative of long term change and climate sensitivity (e.g. from *abrupt-4xCO<sub>2</sub>*  
416 or *amip-p4K*). The experiment has the additional benefit, by differencing with the standard *amip* run that includes time-  
417 varying forcing agents, of providing detailed information on the transient effective radiative forcing and adjustments in  
418 models during the AMIP period (Andrews, 2014). This can then be compared to the forcings diagnosed in the Radiative  
419 Forcing Model Intercomparison Project (RFMIP, Pincus et al., 2016, who use a pre-industrial climate baseline) to test for any  
420 dependence of forcing and adjustments on the climate state. Time-varying feedbacks in the *amip* experiment could  
421 alternatively be diagnosed by subtracting a time-varying radiative forcing diagnosed from RFMIP experiments. However, the  
422 *amip-piForcing* approach has the benefit of diagnosing the time-varying feedbacks over the full 1870-present period rather  
423 than the last 36 years, and does so with reference to a single experiment, which reduces noise compared to that which would  
424 be present with a double difference of the *amip* experiment and two RFMIP experiments. Also, the inclusion of CFMIP  
425 process diagnostics in the *amip-piForcing* experiment will enable a deeper understanding of the factors underlying forcing  
426 and feedback differences in the present and future climate.

427 We also consider the time variation of feedbacks in *abrupt-4xCO<sub>2</sub>* experiments to be an important area to be investigated,  
428 as this can have a substantial impact on estimates of equilibrium sensitivity (e.g. Geoffroy et al., 2013). Andrews et al., 2015  
429 investigated such effects using two atmosphere-only GCMs forced with SSTs and sea ice from their own *abrupt-4xCO<sub>2</sub>*  
430 experiments, and attributed the time variation in the feedbacks to changes in the pattern of surface warming. Pilot studies are  
431 ongoing to develop similar experiments based on a composite SST pattern response more representative of the CMIP5  
432 ensemble mean. We plan to organise an informal pilot intercomparison based on this within CFMIP-3 and may subsequently  
433 propose these experiments as an extension to the CFMIP-3/CMIP6 experiment set.

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## 434 2.7 Time slice experiments for understanding regional climate responses to CO<sub>2</sub> (Tier 2)

435 Lead Coordinators: Robin Chadwick, Hervé Douville and Christopher Skinner

436  
437 Science questions:

- 438 • How do regional climate responses (e.g. of precipitation) in a coupled model arise from the combination of  
439 responses to different aspects of CO<sub>2</sub> forcing and sea surface warming (uniform SST warming, patterned SST  
440 warming, sea-ice change, direct CO<sub>2</sub> effect, plant physiological effect)?

- 442 • Which aspects of forcing/warming are most important for causing inter-model uncertainty in regional climate  
443 projections?
- 444 • Can inter-model differences in regional projections be related to underlying structural or resolution differences  
445 between models through improved process understanding, and could this help us to constrain the range of regional  
446 projections?
- 447 • What impact do coupled model SST biases have on regional climate projections?  
448

449 The CFMIP-2/CMIP5 set of idealised amip experiments (e.g. *amip4K*, *amipFuture*) have allowed the contribution of different  
450 aspects of SST warming and increased CO<sub>2</sub> concentrations to the projections of fully coupled GCMs to be examined (e.g.  
451 Bony et al., 2013; Chadwick et al., 2014; He and Soden, 2015). However the amip experiments were not designed to replicate  
452 coupled GCM responses on a regional scale, and large discrepancies exist between the two in many regions, particularly  
453 when individual models are examined instead of the ensemble mean (Chadwick, 2016). This is largely due to the choice of  
454 present-day and future SST boundary conditions used in the amip experiments, as well as missing processes such as the plant  
455 physiological response to CO<sub>2</sub>, rather than the lack of air-sea coupling (Skinner et al., 2012).

456 We propose a new set of 7 30-year atmosphere-only time slice experiments, and one 36-year amip-style experiment, to  
457 decompose the regional responses of each model's *abrupt-4xCO2* run into separate responses to each aspect of forcing and  
458 warming (uniform SST warming, pattern SST change, sea-ice change, increased CO<sub>2</sub>, plant physiological effect). These are  
459 forced with monthly- and annually-varying monthly mean SSTs and sea ice, which reproduce regional precipitation patterns  
460 more accurately than is possible using climatological SST forcing (Skinner et al., 2012). As well as allowing regional  
461 responses in each individual model to be better understood, this set of experiments should prove especially useful for  
462 understanding the causes of model uncertainty in regional climate change.

463 The experiments are:

- 464 1) *piSST* – An AGCM experiment with monthly- and annually-varying SSTs, sea-ice, atmospheric constituents and any  
465 other necessary boundary conditions (e.g. vegetation if required) taken from a section of each model's own *piControl* run,  
466 using the 30 years of *piControl* that are parallel to years 111-140 of its *abrupt-4xCO2* run. Note that dynamic vegetation (if  
467 included in the model) should not be turned on in any of the *piSST* set of experiments;
- 468 2) *piSST-pxK* – same as *piSST*, but with a global spatially and temporally uniform SST anomaly applied on top of the  
469 monthly- and annually- varying *piSST* SSTs. The magnitude of the uniform increase is taken from each model's global,  
470 climatological annual mean open SST change between *abrupt-4xCO2* and *piControl* (using the mean of years 111-140 of  
471 *abrupt-4xCO2*, and the parallel 30-year section of *piControl*). Sea-ice is unchanged from *piSST* values;
- 472 3) *piSST-4xCO2-rad* – same as *piSST* but CO<sub>2</sub> as seen by the radiation scheme is quadrupled;
- 473 4) *piSST-4xCO2* – same as *piSST* but with CO<sub>2</sub> quadrupled, and this increase is seen by both the radiation scheme and the  
474 plant physiological effect. If a model does not include the plant physiological response to CO<sub>2</sub>, then *piSST-4xCO2* can be  
475 omitted from the set of *piSST* experiments for that model;
- 476 5) *a4SST* – same as *piSST*, but with monthly- and annually-varying SSTs taken from years 111-140 of each model's own  
477 *abrupt-4xCO2* experiment instead of from *piControl* (sea ice is unchanged from *piSST*);
- 478 6) *a4SSTice* – same as *piSST*, but with monthly- and annually-varying SSTs and sea-ice taken from years 111-140 of each  
479 model's own *abrupt-4xCO2* experiment instead of from *piControl*;
- 480 7) *a4SSTice-4xCO2* – same as *piSST*, but with monthly- and annually-varying SSTs and sea-ice taken from years 111-140  
481 of each model's own *abrupt-4xCO2* experiment instead of from *piControl*. CO<sub>2</sub> is also quadrupled, and is seen by both the  
482 radiation scheme and the plant physiological effect (if included in the model). *a4SSTice-4xCO2* is used to establish whether a  
483 time slice experiment can adequately recreate the coupled *abrupt-4xCO2* response in each model, and then forms the basis for  
484 a decomposition using the other experiments. The time slice experiments can be combined in various ways to isolate the  
485 climate response to each individual aspect of forcing and warming. For example the response to SST pattern change is given  
486 by taking the difference between *a4SST* and *piSST-pxK*, and the plant physiological response is found by taking the  
487 difference between *piSST-4xCO2* and *piSST-4xCO2-rad*.
- 488 8) We also propose an additional amip based experiment, *amip-a4SST-4xCO2*: the same as amip, but a patterned SST  
489 anomaly is applied on top of the monthly- and annually-varying amip SSTs. This anomaly is a monthly climatology, taken  
490 from each model's own *abrupt-4xCO2* run minus *piControl* (using the mean of years 111-140 of *abrupt-4xCO2*, and the  
491 parallel 30-year section of *piControl*). CO<sub>2</sub> is quadrupled, and the increase in CO<sub>2</sub> is seen by both the radiation scheme and  
492 vegetation. Comparison of *amip-a4SST-4xCO2* and *a4SSTice-4xCO2* should help to illuminate the impact of SST biases on  
493 regional climate responses in each model, and how this contributes to inter-model uncertainty.

### 494 3 CFMIP Recommended Diagnostic Outputs for CMIP experiments

495 | The CFMIP-3/CMIP6 specific diagnostic request is designed to address the following questions: 1) How well do clouds and  
496 other relevant variables simulated by models agree with observations? 2) What physical processes and mechanisms are  
497 important for a credible simulation of clouds, cloud feedbacks and cloud adjustments in climate models? 4) Which models  
498 have the most credible representations of processes relevant to the simulation of clouds? 5) How do clouds and their changes  
499 interact with other elements of the climate system?

500 | The set of diagnostic outputs recommended for CFMIP-3/CMIP6 is based on that from CFMIP-2/CMIP5, with some  
501 modifications. The request outlined below is in three parts. The first part describes an updated set of CFMIP process  
502 diagnostics (based on those in CFMIP-2/CMIP5 which are documented at [http://cmip-pcmdi.llnl.gov/cmip5/output\\_req.html](http://cmip-pcmdi.llnl.gov/cmip5/output_req.html))  
503 in terms of the various groups of variables and the experiments in which they are requested. This set was drawn up by the

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506 CFMIP committee and ratified by the modelling groups following a presentation at the 2014 CFMIP meeting. The second  
507 part describes recommendations for COSP outputs in the [CFMIP-3/CMIP6](#), CMIP DECK and CMIP6 Historical experiments.  
508 The third part describes additional diagnostics requested for evaluation of mean diurnal cycle of tropical clouds and radiation.  
509 The summaries below give an overview of the diagnostic request; however the definitive and detailed specification is  
510 documented in the CMIP6 data request, available at <https://www.earthsystemcog.org/projects/wip/CMIP6DataRequest>  
511 (Juckes et al., in preparation.) The changes in the [CFMIP-3/CMIP6](#) diagnostics relative to those requested for [CFMIP-](#)  
512 [2/CMIP5](#) are additionally motivated and detailed in the CFMIP CMIP6 proposal document which is available from the  
513 CFMIP website.

514 CMIP mandates that for participation in [CFMIP-3/CMIP6](#), modelling groups must commit to performing all of the Tier 1  
515 experiments. In recognition that sufficient resources are not available for all groups to prepare all of the [CFMIP-3/CMIP6](#)  
516 specific diagnostics, these diagnostics are considered to be Tier 2, i.e., not compulsory for participation in [CFMIP-3/CMIP6](#).  
517 Nonetheless, these diagnostics are extremely valuable and all groups with the capacity to do so are very strongly encouraged  
518 to provide the additionally requested [CFMIP-3/CMIP6](#) specific diagnostics.

519 In the case where [CFMIP-3/CMIP6](#) specific outputs are requested in DECK and CMIP6-Historical experiments, and  
520 modelling groups run more than one ensemble member of an experiment, we request that each set of [CFMIP-3/CMIP6](#)  
521 specific outputs are submitted for one ensemble member only. Having different CFMIP variables in different ensemble  
522 members is acceptable, but submitting them all in the same ensemble member is preferable. We request that the modelling  
523 groups provide information on which CFMIP diagnostic sets are submitted in which ensemble members so that this  
524 information can be made available to those who may be analyzing the output. Our analysis plans for the CFMIP diagnostic  
525 outputs in the CMIP DECK, CMIP6 Historical and [CFMIP-3/CMIP6](#) experiments, including details of the CFMIP  
526 Diagnostics Code Catalogue are summarised in Appendix A.  
527

### 528 3.1 Process outputs

529 In [CFMIP-2/CMIP5](#), instantaneous high frequency 'cfSites' outputs were requested for 120 locations in the *amip*, *amip4K*,  
530 *amipFuture* and *amip4xCO2* experiments, and for 73 locations along the Greenwich meridian in the aquaplanet experiments,  
531 to support understanding and evaluation of clouds and their interactions with convection and other processes. The 120  
532 locations include the locations of instrumented sites (ARM and CloudNet stations, Dome C, etc), the transect associated with  
533 the GCSS Pacific Cross-section Intercomparison (GPCI), past field campaigns (DYCOMS-II, NARVAL, HOPE, VOCALS,  
534 ASTEX and AMMA transects, TOGA-COARE, RICO, etc) and a number of climate regimes that contribute substantially to  
535 the inter-model spread of cloud feedbacks in climate change (Webb et al., 2015a). These outputs have so far been used to  
536 evaluate the models with in-situ measurements (e.g. Nuijens et al., 2015a, Nuijens et al., 2015b, Neggers et al., 2015), to  
537 investigate the diurnal cycle of cloud feedbacks (Webb et al., 2015a) and to compare cloud feedbacks in climate models with  
538 SCM and LES outputs from CGILS (Dal Gesso et al., 2015). We have added St. Helena to the list of locations in light of  
539 upcoming field work, increasing the total number of locations to 121 for [CFMIP-3/CMIP6](#). A text file containing the list of  
540 locations is available in the Supplementary Information and on the CFMIP website; these are also presented graphically in  
541 Figure 2.

542 For CFMIP-3 cfSites outputs are now requested for one ensemble member of the *amip* DECK experiment, and the *amip-*  
543 *p4K* and *amip-4xCO2* experiments. Outputs should be provided for the full duration of each experiment. The sampling  
544 interval should be the integer multiple of the model time step that is nearest to 30 minutes and divides into 60 minutes with  
545 no remainder: e.g. 30 minutes for a 30, 15 or 10 minute time step or 20 minutes for a 20 minute time step. Outputs should be  
546 instantaneous (i.e. not time means) and from nearest grid box (i.e. no spatial interpolation). We have dispensed with the  
547 cfSites outputs in the aquaplanet and *amip-future4K* experiments because these have been less widely used compared to those  
548 from the other experiments.

549 The cfSites outputs from [CFMIP-3/CMIP6](#) provide instantaneous outputs of a range of quantities (including temperature  
550 and humidity tendency terms) in experiments which can be used to evaluate the present day relationships of clouds to cloud  
551 controlling factors using in situ measurements, and at the same time explore how these relationships affect cloud feedbacks  
552 and cloud adjustments. An increasing wealth of observational data with which to evaluate the models using these outputs is  
553 available or in the planning stage, for example from the Barbados Cloud Observatory (Stevens et al., 2015) the ARM  
554 Program (e.g. Wood et al., 2015; Marchand et al., 2015) or within the German national project on high-definition clouds and  
555 precipitation for climate-prediction, HD(CP)<sup>2</sup>, inclusive of its observational prototype experiment (HOPE), and which has  
556 collected observations over Germany following conventions adopted for CMIP (Andrea Lammert, personal communication).

557 CFMIP-2 also requested cloud, temperature and humidity tendency terms from convection, radiation, dynamics etc. in the  
558 *amip*, *amip4K*, *amipFuture* and *amip4xCO2*, *aquaControl*, *aqua4xCO2* and *aqua4K* experiments, as global monthly mean  
559 outputs and high frequency outputs at fixed locations (Bony et al., 2011). Upward and downward radiative fluxes on model  
560 levels were also requested in these experiments, and for instantaneous CO<sub>2</sub> quadrupling in the *amip* experiment only.  
561 Temperature and humidity tendency terms in particular have been shown to be useful for understanding the roles of different  
562 parts of the model physics in cloud feedbacks (e.g. Webb and Lock 2013; Demoto et al., 2013; Sherwood et al., 2014; Brient  
563 et al., 2015) and cloud adjustments (e.g. Kamae and Watanabe 2012; Ogura et al., 2014) as well as in understanding clouds  
564 and circulation in the present climate (e.g. Williams et al., 2013; Oueslati and Bellon, 2013; Xavier et al., 2015). They have  
565 also been used to understand regional warming patterns such as polar amplification in coupled models (e.g. Yoshimori et al.,  
566 2014).

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580 | In [CFMIP-3/CMIP6](#) we have improved the definitions of the temperature and humidity tendency terms, and added some  
581 additional terms such as clear-sky radiative heating rates to more precisely quantify the contributions of different processes to  
582 the temperature and humidity budget changes underlying cloud feedbacks and adjustments. We have dispensed with the cloud  
583 water tendency terms because these have been less widely used than the temperature and humidity tendencies.  
584 A shortcoming of the CMIP5 protocol was that we were unable to interpret the physical feedback mechanisms in coupled  
585 model experiments due to a lack of process diagnostics. For this reason in CMIP6 we are requesting these budget terms in  
586 the DECK *abrupt-4xCO2* experiment and the pre-industrial control as well as one ensemble member of the *amip* DECK  
587 experiment, and all of the [CFMIP-3/CMIP6](#) experiments listed in Sections 2.1-2.6.

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588 Clustering approaches (e.g., Jakob and Tselioudis, 2003) are now commonly used for assessing the contributions of  
589 different cloud regimes (e.g. stratocumulus, trade cumulus, frontal clouds, etc) to present day biases in cloud simulations and  
590 to inter-model differences in cloud feedbacks (e.g. Williams and Webb 2009, Tsushima et al., 2013, Tsushima et al., 2015).  
591 We have also added some additional daily 2D fields to the standard package of CFMIP daily outputs to allow further  
592 investigation of feedbacks between clouds and aerosols associated with the changing hydrological cycle (aerosol loadings and  
593 cloud top effective radii/number concentrations) and a clearer diagnosis of the roles of convective and stratiform clouds  
594 (convective vs. stratiform ice and condensed water paths and cloud top effective radii/number concentrations).

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### 595 3.2 COSP outputs

596 | This section motivates and summarizes the COSP outputs requested from the DECK, and CMIP6 historical and [CFMIP-](#)  
597 [3/CMIP6](#) experiments as well as a corresponding set of observations.

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598 There is no unique definition of clouds or cloud types, neither in models nor in observations. Therefore, to compare models  
599 with observations, and even to compare models with each other, it is necessary to use a consistent definition of clouds  
600 between the model and the satellite product in question (i.e., be “definition-aware”). Further complicating matters - climate  
601 model grid boxes (typically 1 degree) are much larger than the scales over which many satellite observations are made  
602 (typically <10 km). As a result, one must downscale the climate model cloud properties to the observation scale (i.e., be  
603 “scale-aware”). The CFMIP Observation Simulator Package (COSP) enables definition-aware and scale-aware comparisons  
604 between models and multiple sets of observations by producing cloud diagnostics from model simulations that are  
605 quantitatively comparable to a variety of satellite products from ISCCP, CloudSat, CALIPSO, MODIS, MISR and Parasol  
606 (Bodas-Salcedo et al., 2011). COSP enables a more quantitative comparison of model outputs with satellite cloud products,  
607 which often sub-sample low level clouds in the presence of high level clouds due to the effects of cloud overlap and  
608 attenuation (e.g. Yu et al., 1996). COSP also provides histograms of various cloud properties as a function of height or  
609 pressure which are directly comparable with satellite products and cannot be calculated correctly from time mean model  
610 outputs. The multiple simulators within COSP allow a multi-faceted evaluation of clouds in models whereby the strengths  
611 and weaknesses of different satellite products may be considered together.

612 COSP is increasingly being used not only for model intercomparison activities but as part of the model development and  
613 evaluation process by modelling groups (e.g. Marchand et al., 2009; Zhang et al., 2010; Kay et al., 2012; Franklin et al.,  
614 2013; Lacagnina and Selten, 2014; Nam et al., 2014; Williams et al., 2015, Konsta et al., 2015). Many of the standard  
615 monthly and daily COSP outputs have been shown to be valuable in the CMIP5 experiments, not only for cloud evaluation,  
616 allowing a detailed evaluation of clouds and precipitation, and their interaction with radiation (e.g. Nam et al., 2012; Cesana  
617 and Chepfer, 2012; Kay et al. 2012; Klein et al., 2013; Tsushima et al., 2013; Gordon and Klein, 2014; Lin et al., 2014;  
618 Bodas-Salcedo et al., 2014; Bellomo and Clement, 2015), but also in quantifying the contributions of different cloud types to  
619 cloud feedbacks and forcing adjustments in climate change experiments (e.g. Zelinka et al., 2013; Zelinka et al., 2014;  
620 Chepfer et al., 2014; Tsushima et al., 2015). For a full list of studies that use COSP diagnostics for model evaluation and  
621 feedback analysis please refer to the ‘CFMIP publications’ section of the CFMIP website.

622 Here we will give only a brief overview of the COSP request; readers interested in the complete details of the data request  
623 are referred to the Earth System CoG website (<https://www.earthsystemcog.org/projects/wip/CMIP6DataRequest>). The  
624 COSP data request for the CMIP DECK and CMIP6 has been designed to span model evaluation across different space and  
625 time scales. Monthly-mean diagnostics allow for the evaluation and intercomparison of large-scale distributions of cloud  
626 properties and their interaction with radiation. High-frequency model outputs (daily, 3-hourly) are aimed at a process-oriented  
627 evaluation (e.g. Bodas-Salcedo et al., 2012) and offer the opportunity of exploiting the synergy between multiple instruments  
628 (e.g. Konsta et al., 2015). Recent observational developments have improved our capability to retrieve cloud radiative  
629 properties. In particular, new methodologies for cloud phase identification are available for CALIPSO and MODIS, and  
630 COSP has been enhanced to provide diagnostics that are compatible with these new observational datasets (Cesana and  
631 Chepfer, 2013). These new diagnostics will help elucidate some open questions regarding the role of cloud phase in model  
632 biases (Ceppi et al., 2016; Bodas-Salcedo et al., in press).

634 | Within [CFMIP-3/CMIP6](#), COSP output is requested from six simulators as follows:

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- 635 • ISCCP: pseudo-retrievals of cloud top pressure (CTP) and cloud optical thickness (tau) (Klein and Jakob 1999;  
636 Webb et al., 2001).
- 637 • CloudSat: a forward model for radar reflectivity as a function of height (Haynes et al., 2007).
- 638 • CALIPSO (Chepfer et al., 2008; Cesana and Chepfer, 2013): forward model for lidar scattering ratio as function of  
639 height, and cloud phase retrieval.
- 640 • MODIS: pseudo-retrievals of CTP, effective particle size and tau as function of phase (Pincus et al., 2012).

- MISR: pseudo-retrievals of cloud top height (CTH) and tau (Marchand and Ackerman, 2010).
- PARASOL: simple forward model of mono-directional reflectance (Konsta et al., 2015).

The main difference to CFMIP-2 is that output is requested from a greater number of simulators and longer periods of simulated time. MISR provides more accurate retrievals of cloud-top-height for low-level and mid-level clouds, and more reliable discrimination of mid-level clouds from other clouds, while MODIS provides better retrievals of high-level clouds. ISCCP and MISR histograms can be combined to separate optically-thin high-level clouds into multi-layer and single-layer categories (Marchand et al. 2010). Aerosol schemes are becoming more complex, with more elaborate representations of cloud-aerosol interactions. This makes the evaluation of the phase partitioning an important aspect of model evaluation, and height-resolved partitioning estimates from the CALIPSO simulator are included in the COSP request. Cloud phase and particle size estimates from the MODIS simulator were not available in CFMIP-2 but may prove a useful complement to investigate cloud-aerosol interactions by virtue of greater geographic sampling and longer time records. Many of the COSP diagnostics are now requested for the entire lengths of the DECK, CMIP6 Historical and [CFMIP-3/CMIP6](#) experiments to support the quantification and interpretation of cloud feedbacks and cloud adjustments in a broader context. The new inclusion in this COSP request of a long time series of three-dimensional cloud fractions will facilitate the comparison of cloud trends with the observational record (Chepfer et al., 2014). More details of all the changes with respect to [CFMIP-2/CMIP6](#) can be found in the proposal of the CMIP6-Endorsed MIPs, available from the CMIP6 website (<http://www.wcrp-climate.org/wgcm-cmip/wgcm-cmip6>).

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The COSP output is in six variable groups:

- 1) cfMon\_sim: monthly means of ISCCP 2D diagnostics (cloud fraction, cloud albedo, and cloud top pressure), ISCCP CTP-tau histogram, and CALIPSO 2D and 3D cloud fractions.
- 2) cfDay\_2d: daily means of ISCCP and CALIPSO 2D diagnostics, and PARASOL reflectances.
- 3) cfDay\_3d: daily means of ISCCP and CALIPSO 3D diagnostics.
- 4) cfMonExtra: monthly means of CloudSat reflectivity and CALIPSO scattering ratio histograms as function of height, CALIPSO 3D cloud fractions by phase, MODIS 2D cloud fractions, MODIS CTP-tau histogram and size-tau histograms by phase, MISR CTH-tau histograms, and PARASOL reflectances.
- 5) cfDayExtra: daily means of CALIPSO total cloud fraction, MODIS CTP-tau histogram and size-tau histograms by phase, and PARASOL reflectances.
- 6) cf3hrSim: 3-hourly instantaneous diagnostics of ISCCP CTP-tau histograms, MISR CTH-tau histograms, MODIS CTP-tau histogram and size-tau histograms by phase, CALIPSO 2D and 3D cloud fractions, CloudSat reflectivity and CALIPSO scattering ratio histograms as function of height, and PARASOL reflectances.

The variable groups cfMon\_sim and cfDay\_2d are requested for all years in the *amip* experiment performed as part of the DECK and the CMIP6-Historical experiments, and for 140 years of the *piControl*, *IpctCO2*, and *abrupt-4xCO2*. These are requested for one ensemble member only from these experiments. They are also requested in all of the [CFMIP-3/CMIP6](#) experiments listed in Sections 2.1-2.6 above. cfDay\_3d is requested in one ensemble member of the DECK *amip* experiment and in the [CFMIP-3/CMIP6](#) *amip-p4K* and *amip-4xCO2* experiments. cfMonExtra and cfDayExtra are requested for all years of one ensemble member of the *amip* DECK experiment, and cf3hrSim for the year 2008 only. (Please note that in the full data request these variable groups are in many cases split into a number of sub-tables. As noted above, the formal data request provides the definitive specification of the model outputs.)

COSP is available via the CFMIP website (<https://www.earthsystemcog.org/projects/cfmip>). Version 1.4 is a stable code release that was made available well in advance of CMIP6 at the request of the modelling groups. Small updates are required to enable some new diagnostics requested by CFMIP-3/CMIP6, most notably joint histograms of particle size and optical thickness from the MODIS simulator; with these updates the code is known as version 1.4.1. Modeling centres are encouraged to update to COSP 1.4.1 to provide these new diagnostics but may provide results from COSP 1.4.

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Developed over the last few years, COSP 2 substantially revises the infrastructure for integrating satellite simulators in climate models. COSP 2 makes many fewer inherent assumptions about the model representation of clouds than do previous versions but contains an optional interface allowing it to be used as a drop-in replacement for COSP 1.4 or COSP 1.4.1. At the time of this writing COSP 2 is undergoing final testing in two climate models. Availability of the final version will be announced on the CFMIP website and modelling groups are free to adopt it for use in CFMIP at that time.

The CFMIP community has developed a set of observational datasets available via the CFMIP-OBS website (<http://climserv.ipsl.polytechnique.fr/cfmip-obs/>) that are defined consistently with the COSP diagnostics and the [CFMIP-3/CMIP6](#) data request in terms of vertical grids and time averaging periods. These are mostly reported as monthly means although some are reported at higher temporal resolution for process oriented model evaluations (e.g. Konsta et al., 2012). Table 3 summarizes the datasets relevant to the COSP CMIP6 data request. Some of the CFMIP-OBS datasets listed in Table 3 (CALIPSO, CloudSat, ISCCP, PARASOL) are also available from the ESGF as part of the obs4MIPs project (Teixeira et al., 2014). These datasets are periodically updated to include more recent data from the relevant satellites, many of which are still operational. Please refer to the CFMIP-OBS website for updates.

### 3.3 Monthly Mean Diurnal Cycle Outputs

711 Climate models have difficulties representing the diurnal cycle of convective clouds over land (Yang and Slingo, 2001;  
712 Stratton and Stirling, 2011), but its evaluation is not possible with sun-synchronous satellites. Geostationary satellites  
713 provide high-frequency sampling that can be used to evaluate model biases in the diurnal cycle of clouds and radiation (albeit  
714 over a limited area). The Geostationary Earth Radiation Budget instrument (GERB; Harries et al., 2005) measures the top of  
715 atmosphere (TOA) radiation budget from a geostationary orbit at 0E at 15 minute frequency, which provides a unique view of  
716 tropical convection over Africa. The variable group *cf1hrClimMon* requests monthly mean diurnal cycle of TOA radiative  
717 fluxes (all-sky and clear sky) for the entire length of the *amip* DECK experiment. The radiative fluxes are hourly UTC means.  
718 The ‘average day’ for each month of the simulation is then constructed by averaging each UTC hourly mean over the entire  
719 month. These diagnostics will be directly comparable with GERB measurements.  
720

#### 721 4. Summary

722 The primary goal of CFMIP is to inform improved assessments of cloud feedbacks on climate change. This involves bringing  
723 climate modelling, observational and process modelling communities closer together and providing better tools and  
724 community support for understanding and evaluation of clouds and cloud feedbacks simulated by climate models. CFMIP  
725 supports ongoing coordinated model inter-comparison activities by recommending experiments and model output diagnostics  
726 for CMIP, designed to support the understanding and evaluation of cloud processes and cloud feedbacks in models. The  
727 CFMIP approach is also increasingly being used to understand other aspects of climate change, and so a second objective has  
728 now been introduced, to improve understanding of circulation, regional-scale precipitation, and non-linear changes. CFMIP  
729 proposes a number of [CFMIP-3/CMIP6](#) experiments and model outputs for CMIP6, building on and extending those which  
730 were part of [CFMIP-2/CMIP5](#).

731 A compact set of [CFMIP-3/CMIP6](#) Tier 1 experiments are proposed address the question: “1) What are the physical  
732 mechanisms underlying the range of cloud feedbacks and cloud adjustments predicted by climate models, and which models  
733 have the most credible cloud feedbacks?” The Tier 1 experiments (*amip-p4K*, *amip-4xCO2*, *amip-future4K*, *aqua-control*,  
734 *aqua-4xCO2* and *aqua-p4K*) retain the idealized experimental hierarchy of the [CFMIP-2/CMIP5](#) experiments while building  
735 on the DECK AMIP experiment. A number of Tier 2 experiments are proposed to address additional science questions. An  
736 *amip* uniform minus 4K experiment is proposed to address the question “2) Are cloud feedbacks consistent for climate  
737 cooling and warming, and if not, why?” Atmosphere-only experiments with clouds made transparent to longwave radiation  
738 address the question “3) How do cloud-radiative effects impact the structure, the strength and the variability of the general  
739 atmospheric circulation in present and future climates?” Abrupt +/-4% Solar Forced AOGCM experiments are proposed for  
740 the question “4) How do responses in the climate system due to changes in solar forcing differ from changes due to CO<sub>2</sub>, and  
741 is the response sensitive to the sign of the solar forcing?” abrupt 2xCO<sub>2</sub> and abrupt 0.5xCO<sub>2</sub> experiments are proposed to  
742 address the question “5) To what extent is regional-scale climate change per CO<sub>2</sub> doubling state-dependent (nonlinear), and  
743 why?” Other experiments and questions proposed include: AMIP with preindustrial forcing “6) Are climate feedbacks during  
744 the 20<sup>th</sup> century different to those acting on long term climate change and climate sensitivity?”; Time slice experiments forced  
745 with SSTs from preindustrial and *abrupt-4xCO2* simulations “7) How do regional climate responses (of e.g. precipitation) in  
746 a coupled model arise from the combination of responses to different aspects of CO<sub>2</sub> forcing and warming (uniform SST  
747 warming, pattern SST warming, direct CO<sub>2</sub> effect, plant physiological effect, sea-ice change)?”

748 The [CFMIP-3/CMIP6](#) experiments will continue to include outputs from the CFMIP Observational Simulator Package  
749 (COSP) to support robust scale-aware and definition-aware evaluation of modelled clouds with observations and to relate  
750 cloud feedbacks to observed quantities. COSP outputs are also proposed for inclusion in the DECK and CMIP6 Historical  
751 experiments. Process diagnostics including ‘cfSites’ high frequency outputs at selected locations and temperature and  
752 humidity budget terms from radiation, convection, dynamics, etc. are also retained from [CFMIP-2/CMIP5](#). These will help to  
753 address the following questions: 1) How well do clouds and other relevant variables simulated by models agree with  
754 observations? 2) What physical processes and mechanisms are important for a credible simulation of clouds, cloud feedbacks  
755 and cloud adjustments in climate models? 4) Which models have the most credible representations of processes relevant to  
756 the simulation of clouds? 5) How do clouds and their changes interact with other elements of the climate system?  
757 By continuing the [CFMIP-2/CMIP5](#) experiments and diagnostic outputs within [CFMIP-3/CMIP6](#) we hope to apply the well  
758 established aspects of the CFMIP approach to a larger number of climate models. Additionally we have proposed new  
759 [CFMIP-3/CMIP6](#) experiments to investigate a broader range of questions relating to the Grand Challenge on Clouds,  
760 Circulation and Climate Sensitivity. We hope that the modelling community will participate fully in [CFMIP-3](#) via CMIP6 so  
761 as to maximize the relevance of our findings to future assessments of climate change.

#### 762 Code and Data Availability

763 COSP is published under an open source license via GitHub (please see the CFMIP website for details). The model output  
764 from the DECK, CMIP6 historical and [CFMIP-3/CMIP6](#) simulations described in this paper will be distributed through the  
765 Earth System Grid Federation (ESGF) with digital object identifiers (DOIs) assigned. As in CMIP5, the model output will be  
766 freely accessible through data portals after registration. In order to document CMIP6’s scientific impact and enable ongoing  
767 support of CMIP, users are obligated to acknowledge CMIP6, the participating modelling groups, and the ESGF centres (see  
768 details on the CMIP Panel website at <http://www.wcrp-climate.org/index.php/wgcm-cmip/about-cmip>). Further information  
769 about the infrastructure supporting CMIP6, the metadata describing the model output, and the terms governing its use are  
770 provided by the WGCM Infrastructure Panel (WIP) in their invited contribution to this Special Issue. Along with the data

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774 itself, the provenance of the data will be recorded, and DOIs will be assigned to collections of output so that they can be  
775 appropriately cited. This information will be made readily available so that published research results can be verified and  
776 credit can be given to the modelling groups providing the data. The WIP is coordinating and encouraging the development of  
777 the infrastructure needed to archive and deliver this information. In order to run the experiments, datasets for natural and  
778 anthropogenic forcings are required. These forcing datasets are described in separate invited contributions to this Special  
779 Issue. The forcing datasets will be made available through the ESGF with version control and DOIs assigned.

780  
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## 794 Appendix A: Analysis Plan and CFMIP Diagnostic Codes Catalogue

795 CFMIP-2 analysis activities are ongoing and the CFMIP community is ready to analyse CFMIP-3/CMIP6 data at any time.  
796 We would like modelling groups to perform the proposed CFMIP-3/CMIP6 experiments at the same time or shortly after  
797 their DECK and CMIP6 Historical experiments. Subsequent informally organised CFMIP-3 experiments which are not  
798 included in CMIP6 will build on the proposed DECK and CFMIP-3/CMIP6 experiments and some will start as soon as  
799 CMIP6 DECK experiments start to become available. We envisage a succession of CFMIP related intercomparisons  
800 addressing different questions arising from the Grand Challenge spanning the duration of CMIP6.

801 We plan to scientifically analyze, evaluate and exploit the proposed experiments and diagnostic outputs, and have  
802 identified lead coordinators within CFMIP for different aspects of this activity. The lead coordinators are responsible for  
803 encouraging analysis of the relevant experiments as broadly as possible across the scientific community. While they may lead  
804 some analysis themselves, they do not have any first claim on analysing or publishing the results. All interested investigators  
805 are encouraged to exploit the data from these experiments. While investigators may wish to liaise with the lead coordinators  
806 to avoid duplicating work that others are doing, this is not a requirement. An overview of the proposed evaluation/analysis of  
807 the CMIP DECK, CMIP6 Historical and CFMIP-3/CMIP6 experiments follows:

808 CFMIP will continue to exploit the CMIP DECK and CMIP6 experiments to understand and evaluate cloud processes and  
809 cloud feedbacks in climate models. The wide range of analysis activities described above in the context of CFMIP-2 will be  
810 continued in CFMIP-3 using the CMIP DECK and CFMIP-3/CMIP6 experiments, allowing the techniques developed in  
811 CFMIP-2 to be applied to an expanding number of models, including the new generation of models currently under  
812 development. These activities will include evaluation of clouds using additional simulators, investigation of cloud processes  
813 and cloud feedback/adjustment mechanisms using process outputs (cfSites, tendency terms, etc). The inclusion of COSP and  
814 budget tendency terms in additional DECK experiments (e.g. *abrupt-4xCO2*) will enable the CFMIP approach to be applied  
815 to a wider range of experimental configurations. Lead coordinator: Mark Webb.

816 Analysis of the +/-4% solar forcing runs will include an evaluation of both rapid adjustments and longer-term responses on  
817 global and regional top-of-atmosphere radiative fluxes, cloud types (using ISCCP and other COSP simulators) and  
818 precipitation characteristics, as well as comparison of these responses with responses in DECK *abrupt-4xCO2* experiments.  
819 GeoMIP and SolarMIP have expressed a strong interest in these CFMIP-3/CMIP6 experiments and joint analysis of these  
820 experiments with GeoMIP and SolarMIP experiments is anticipated, specifically with the goal of determining to what degree  
821 results from abrupt solar forcing only experiments and abrupt CO<sub>2</sub> only experiments can be used to predict what happens  
822 when both forcing are applied simultaneously, as done in the GeoMIP experiments. Lead coordinators: Chris Bretherton,  
823 Roger Marchand and Bjorn Stevens.

824 Analysis of nonlinear climate processes is discussed in detail by Good et al., 2016. This includes a method for validating  
825 traceability of abrupt CO<sub>2</sub> experiments to transient simulations, which is also recommended as a standard test of the DECK  
826 *abrupt-4xCO2* experiment. Analysis will primarily involve comparing the *abrupt-4xCO2*, *abrupt-2xCO2* and *abrupt-*  
827 *0p5xCO2* experiments over the same timescale. Lead coordinator: Peter Good.

828 Analysis of *amip-piForcing* has already been performed in detail for two models in Andrews, 2014 and Gregory and  
829 Andrews (submitted). We propose to use this as a starting point for a multi-model analysis. Lead coordinator: Timothy  
830 Andrews.

831 An overview analysis of regional responses and model uncertainty in the piSST set of experiments will be carried out by  
832 the coordinators, in collaboration with members of contributing modelling groups. We anticipate that further detailed analysis  
833 on the processes at work in different regions will be carried out by a variety of research groups with interest and expertise in a  
834 particular region: for example a set of similar experiments has previously been used to examine the climate response of the  
835 West African monsoon in CCSM3 (Skinner et al., 2012). The piSST set of experiments have already been successfully run

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841 using the Met Office, NCAR and CNRM CMIP5 models. Lead Coordinators: Robin Chadwick, Hervé Douville and  
842 Christopher Skinner.

843 The analysis of the COOKIE experiments will be reviewed by the coordinators in collaboration with members of the  
844 contributing modelling groups. The role of longwave atmospheric cloud-radiative effects in large-scale circulations, regional  
845 precipitation patterns and the organisation of tropical convection will be investigated in the current climate and in climate  
846 change, with the aim of highlighting both robust effects and sources of uncertainties in the model responses. Lead  
847 coordinators: Sandrine Bony and Bjorn Stevens.

848 When analyzed together with the *amip-p4K* experiment, the *amip-m4K* experiment allows the CFMIP process diagnostics  
849 to be used to understand for asymmetries in the climate response to warming and cooling which have been noted in PMIP  
850 experiments. These might arise from cloud phase responses in middle- and high-latitude clouds or from the adiabatic cloud  
851 liquid water path response feedback which is important over land regions and which would be expected to be weaker with  
852 cooling because of the non-linearity in the Clausius-Clapeyron relation. Lead coordinators: Mark Webb and Bjorn Stevens.

853 The COSP data request for the *amip* DECK experiment will allow a comprehensive multi-model evaluation of clouds and  
854 radiation, following on from CMIP5 studies (e.g. Klein et al., 2013; Bodas-Salcedo et al., 2014). The COSP data request for  
855 the other experiments (e.g. *amip-p4K*, *abrupt-4xCO2*, etc.) permits evaluation of cloud feedbacks and adjustments by cloud  
856 type (Zelinka et al., 2013, Tsushima et al., 2015) or cloud trends (Chepfer et al., 2014). New COSP diagnostics have been  
857 used in single-model analyses: cloud phase diagnostics (Cesana and Chepfer, 2013); MISR simulator outputs to evaluate  
858 cloud fraction and multilayer clouds (Marchand and Ackerman, 2010); CALIPSO vertical distribution of cloud fraction for  
859 the study of cloud trends (Chepfer et al., 2014). These studies will be used as starting points for multi-model analyses. The  
860 COSP Project Management Committee co-chairs will coordinate and encourage the exploitation of these resources. Lead  
861 coordinators: Alejandro Bodas-Salcedo and Steve Klein.

862 Analysis of output from [the CFMIP-3/CMIP6](#) and [CMIP DECK](#) experiments will also be facilitated by sharing of  
863 diagnostic codes via the CFMIP Diagnostics Code Catalogue (accessible via the CFMIP website  
864 <http://www.earthsystemcog.org/projects/cfmip/>). This is a catalogue of programs written by various members of the CFMIP  
865 community, implementing a number of diagnostic approaches from published studies. These include daily cloud clustering  
866 evaluation metrics based on ISCCP and ISCCP simulator outputs (Williams and Webb, 2009, Tsushima et al., 2013), error  
867 metrics for total cloud amount, longwave and shortwave cloud properties (Klein et al., 2013), process oriented evaluation of  
868 clouds using A-train instantaneous observations (Konsta et al., 2012), quality control and low-cloud diagnostics (Nam et al.,  
869 2012; Nam and Quaas, 2012), sensitivity of low cloud cover to estimated inversion strength and SST (Qu et al., 2014) and  
870 cloud radiative kernels (Zelinka et al., 2012). Any codes which implement diagnostics which are relevant to analysing  
871 clouds, circulation and climate sensitivity in models and which are documented in peer reviewed studies are eligible for  
872 inclusion in the catalogue, and we welcome additional contributions to further support community analysis of CMIP6  
873 outputs.

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## 874 APPENDIX B: Aquaplanet Experimental Design

875 Aquaplanets are Earth-like planets with completely water-covered surfaces. They are often used as idealized configurations  
876 of atmospheric GCMs, and in this context the usual convention is that landmasses and topography are removed. Although  
877 many flavours of aquaplanet configurations exist, another convention is to retain as much of the atmospheric model's  
878 formulation as possible. That is, the numerical grid, dynamical core, and parameterized physics are all used just as in realistic  
879 climate simulations.

881 The Tier 1 aquaplanet experiments follow the same experimental design as [CFMIP-2/CMIP5](#) (Medeiros et al., 2015). Those,  
882 in turn, were closely related to previous aquaplanet descriptions. In particular, the control configuration closely follows the  
883 AquaPlanet Experiment protocol (Blackburn and Hoskins, 2013) using a prescribed SST pattern described by Neale and  
884 Hoskins (2000). Two additional runs paralleled the [CFMIP-2/CMIP5](#) *amip4K* and *amip4xCO2* experiments: a uniform 4K  
885 warming and a quadrupling of atmospheric CO<sub>2</sub>.

886 Here we provide the detailed experimental protocol for the three aquaplanet simulations that are part of Tier 1. We note  
887 again that these follow the APE protocol and [CFMIP-2/CMIP5](#), and therefore largely mirror previous descriptions in  
888 Blackburn and Hoskins (2013), Williamson et al. (2012), and Medeiros et al. (2015).

889 Orbital parameters are set to perpetual equinox conditions. This is usually achieved by setting eccentricity and obliquity to  
890 zero to define a circular orbit and insolation independent of calendar. The diurnal cycle is retained. Insolation is based on a  
891 non-varying solar constant of 1365 W m<sup>-2</sup>.

892 The SST is non-varying and zonally uniform. The longitudinal variation is specified using the “Qobs” SST pattern from  
893 Neale and Hoskins (2000), given by:

$$894 T(\varphi) = \begin{cases} \frac{1}{2} (2 - \sin^4 \phi - \sin^2 \phi) \delta T + T_{\min}, & \text{if } |\varphi| < \frac{\pi}{3} \\ 0, & \text{otherwise} \end{cases} \quad (B1)$$

895 where  $\varphi$  is latitude,  $\phi = \frac{\pi}{2} \frac{\varphi}{\varphi_{\max}}$ ,  $\varphi_{\max} = \frac{\pi}{3}$ ,  $\delta T = T_{\max} - T_{\min}$ ,  $T_{\max} = 27^\circ\text{C}$ , and  $T_{\min} = 0^\circ\text{C}$ .

896  
897 Because results are sensitive to the specification of the SSTs, groups that use a prognostic equation for the surface skin  
898 temperature are asked to set this skin temperature to the specified SST. No sea ice is prescribed, so the surface temperature is  
899 spatially uniform at 0°C poleward of 60° for the control simulation.

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904 Radiatively active trace gases are well-mixed with mixing ratios following the AMIP II recommendations: CO<sub>2</sub>: 348 ppmv;  
905 CH<sub>4</sub>: 1650 ppbv; N<sub>2</sub>O: 306 ppbv; Halocarbon yield of approximately 0.24 W m<sup>-2</sup> radiative forcing. The ozone distribution is  
906 the same as used in APE and CFMIP-2/CMIP5, and is derived from the climatology used in AMIP II (Gates et al., 1999), and  
907 is constant in time and symmetric zonally and about the equator. This ozone distribution is provided as a netCDF file which is  
908 archived on the Earth System Grid and available via the DOI <http://dx.doi.org/10.5065/D61834Q6>. Ozone values are  
909 provided up to 0.28hPa (about 60km altitude in mid-latitudes). For models with tops above this level, a high top ozone  
910 dataset is also provided, which is available via the DOI <http://doi.org/10.5065/D64X5653>. The ozone climatologies provided  
911 uses pressure as a vertical coordinate. Most models use a sigma or hybrid vertical coordinate in pressure or altitude, which  
912 will mean that the pressure on a given model level varies in time, near the surface at the very least. Although the ozone  
913 climatology can be interpolated to the pressure of each model level as it varies in time within the model, for simplicity we  
914 recommend interpolating the ozone dataset onto the model vertical grid before the experiment is performed, and then  
915 specifying ozone values which are constant in time on each model level. This vertical interpolation will require a zonally  
916 symmetric climatology of pressure on model levels which is as consistent as possible with that expected in the aqua-control  
917 experiment. This could for example be produced by initially running a test version of the aqua-control experiment with an  
918 ozone climatology taken from a more realistic model configuration such as the AMIP DECK experiment.

919 Aerosols are removed to the extent possible to remove aerosol-radiation interaction (aka direct effects) and aerosol-cloud  
920 interaction (aka indirect effects). No external surface emissions are to be prescribed. Models requiring aerosol for cloud  
921 condensation should use a constant oceanic climatology that is symmetric about the equator and zonally. Alternatively,  
922 models with the capability should set the cloud droplet and crystal numbers to 100\*10<sup>6</sup> m<sup>-3</sup> and 0.1\*10<sup>6</sup> m<sup>-3</sup>, respectively (as  
923 in Medeiros et al., 2016).

924 As in APE, it is recommended that the atmospheric dry mass be adjusted to yield a global mean of 101080 Pa. It is also  
925 recommended to adopt the APE recommended values for geophysical constants, as listed in Table 2 of Williamson et al.  
926 (2012).

927 The aqua-4K experiment follows the above protocol, but with SST derived by adding 4K to Eq. B1.

928 The aqua-4xCO<sub>2</sub> experiment replaces the CO<sub>2</sub> mixing ratio with 1392 ppmv. The SST is unchanged from the control  
929 simulation (Eq. B1).

930 Model runs should be 10 years. We recommend discarding the initial spin up period of a few months.

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## 933 | **APPENDIX C: SST Pattern for CFMIP-3/CMIP6 *amip-future4K/amipFuture*** 934 **experiments**

935

936 The *amip-future4K* (formerly *amipFuture*) experiment is the same as the *amip* DECK experiment, except that the SSTs are  
937 subject to a composite SST warming pattern derived from the CMIP3 coupled models. The patterned SST forcing dataset is  
938 available in a netcdf file called *cfmip2\_4k\_patterned\_sst\_forcing.vn1.0.nc* which is available in the supplementary  
939 information for this paper, and via the CFMIP website. This is a normalised multi-model ensemble mean of the ocean  
940 surface temperature response pattern (the change in ocean surface temperature (TOS) between years 0-20 and 140-160, the  
941 time of CO<sub>2</sub> quadrupling in the 1% runs) from thirteen CMIP3 AOGCMs (cccma, cnrm, gfdlcm20, gfdlcm21, giss, er,  
942 inmcm3, ipsl, miroc-medres, miub, mpi, mri, ncar-ccsm3, and ncar-pcm1.) Before computing the multi-model ensemble  
943 mean, each model's TOS response was divided by its global mean and multiplied by 4. This guarantees that the pattern  
944 information from all models is weighted equally and the global mean SST forcing is the same as in the uniform +4K  
945 experiment. We have retained the SST forcing based on the CMIP3 coupled models because we consider it more important to  
946 be able to compare CMIP5 and CMIP6 models forced with the same SST pattern than to use a pattern which is consistent  
947 with, say, the CMIP5 coupled response.  
948

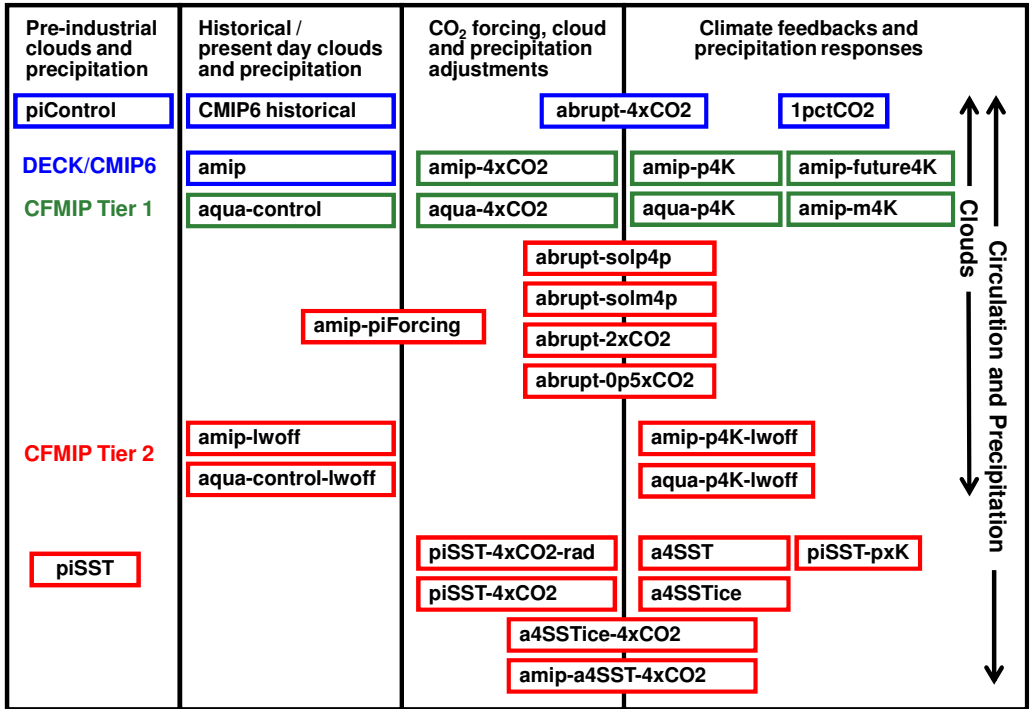
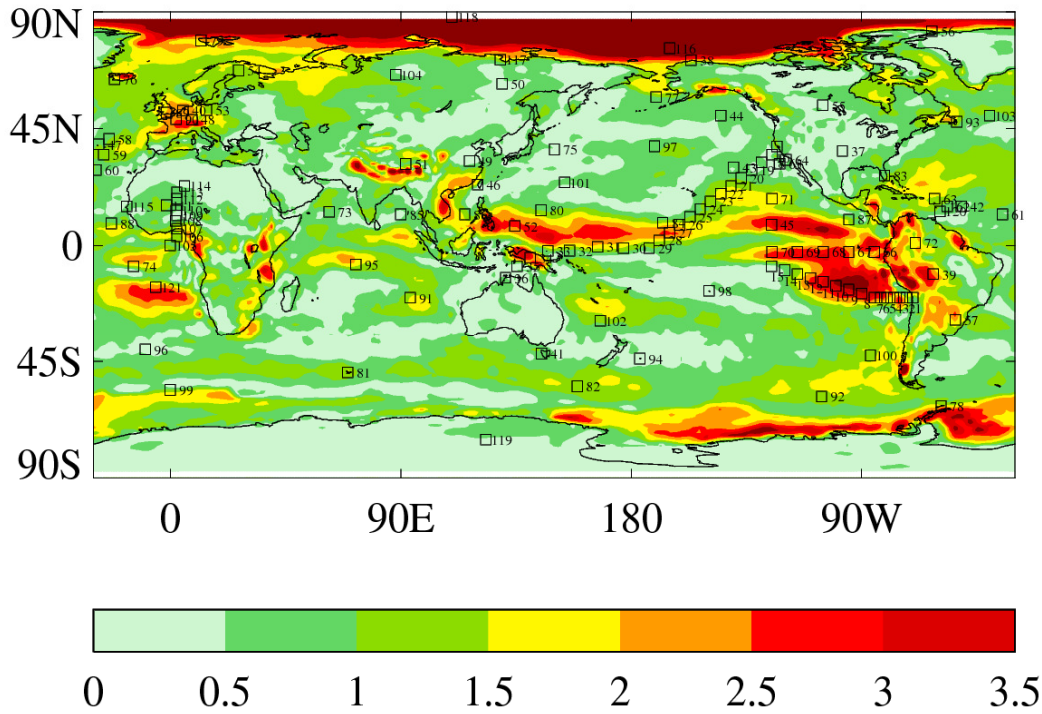


Figure 1. Summary of CFMIP-3/CMIP6 experiments and DECK + CMIP6 Historical experiments.

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## CFMIP-3 cfSites Locations



956 **Figure 2.** ~~CFMIP-3/CMIP6~~ cfSites locations. The contours give an indication of inter-model spread in cloud feedback from  
 957 the ~~CFMIP-2/CMIP5~~ amip/amip4K experiments (please refer to Webb et al., 2015a for details).  
 958  
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962 | **Table 1.** Summary of CFMIP-3/CMIP6 Tier 1 experiments.

Experiment Name	Experiment Description / Design	Configuration	Start Year	Length
amip	This is a single ensemble member of the AMIP DECK experiment which contains additional outputs which are required for model evaluation using COSP, and as control values for model outputs in the amip-p4K, amip-4xCO2, amip-future4K and amip-m4K experiments.	Atmos-only	1979	36
amip-p4K	As CFMIP-2/CMIP5 amip4K experiment. AMIP experiment where SSTs are subject to a uniform warming of 4K.	Atmos-only	1979	36
amip-4xCO2	As CFMIP-2/CMIP5 amip4xCO2 experiment. AMIP experiment where SSTs are held at control values and the CO2 seen by the radiation scheme is quadrupled.	Atmos-only	1979	36
amip-future4K	As CFMIP-2/CMIP5 amipFuture experiment. AMIP experiment where SSTs are subject to a composite SST warming pattern derived from coupled models, scaled to an ice-free ocean mean of 4K.	Atmos-only	1979	36
aqua-control	Extended version of CFMIP-2/CMIP5 aquaControl experiment. Aquaplanet (no land) experiment with no seasonal cycle forced with specified zonally symmetric SSTs.	Atmos-only	1979	10
aqua-p4K	Extended version of CFMIP-2/CMIP5 aqua4K experiment. Aquaplanet experiment where SSTs are subject to a uniform warming of 4K.	Atmos-only	1979	10
aqua-4xCO2	Extended version of CFMIP-2/CMIP5 aqua4xCO2 experiment. Aquaplanet experiment where SSTs are held at control values and the CO2 seen by the radiation scheme is quadrupled.	Atmos-only	1979	10

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965 | **Table 2.** Summary of CFMIP-3/CMIP6 Tier 2 experiments.

Experiment Name	Experiment Description / Design	Configuration	Start Year	Length
amip-m4K	As amip experiment but SSTs are subject to a uniform cooling of 4K.	Atmos-only	1979	36
amip-lwoff	As amip experiment, but with cloud-radiative effects switched off in the LW radiation code.	Atmos-only	1979	36
amip-p4K-lwoff	As amip-p4K experiment, but with cloud-radiative effects switched off in the LW radiation code.	Atmos-only	1979	36
aqua-control-lwoff	As aqua-control experiment, but with cloud-radiative effects switched off in the LW radiation code.	Atmos-only	1979	10
aqua-p4K-lwoff	As aqua-p4K experiment, but with cloud-radiative effects switched off in the LW radiation code.	Atmos-only	1979	10
abrupt-solp4p	Conceptually similar to abrupt 4xCO2 DECK experiment, except that the solar constant rather than CO2 is abruptly increased by 4%.	Coupled AOGCM	1850	150
abrupt-solm4p	Same as abrupt-solp4p, except solar constant is reduced by 4% rather than increased.	Coupled AOGCM	1850	150
abrupt-2xCO2	Identical to the DECK abrupt4xCO2, but at 2xCO2.	Coupled AOGCM	1850	150
abrupt-0p5xCO2	Identical to the DECK abrupt4xCO2, but at 0.5xCO2	Coupled AOGCM	1850	150
amip-piForcing	Identical to AMIP DECK experiment but from 1870-present with constant pre-industrial forcing levels (anthro & natural).	Atmos-only	1870	145
piSST	An AGCM experiment with monthly-varying SSTs, sea-ice, atmospheric constituents and any other necessary boundary conditions (e.g. vegetation if required) taken from each model's own piControl run (using the 30 years of piControl that are parallel to years 111-140 of its abrupt4xCO2 run). Dynamic vegetation should be turned off in all the piSST set of experiments.	Atmos-only	Year 111 of abrupt-4xCO2	30
piSST-pxK	Same as piSST, but with a spatially and temporally uniform SST anomaly applied on top of the monthly-varying piSST SSTs. The magnitude of the uniform increase is taken from each model's global, climatological annual mean open SST change between abrupt4xCO2 minus piControl (using the mean of years 111-140 of abrupt4xCO2, and the parallel 30-year section of piControl).	Atmos-only	Year 111 of abrupt-4xCO2	30
piSST-4xCO2-rad	Same as piSST but CO2 as seen by the radiation scheme is quadrupled.	Atmos-only	Year 111 of abrupt-4xCO2	30
piSST-4xCO2	Same as piSST but CO2 is quadrupled. The increase in CO2 is seen by both the radiation scheme and vegetation.	Atmos-only	Year 111 of abrupt-4xCO2	30
a4SST	As piSST, but with monthly-varying SSTs taken from years 111-140 of each model's own abrupt4xCO2 experiment instead of from piControl. Sea-ice is unchanged from piSST.	Atmos-only	Year 111 of abrupt-4xCO2	30
a4SSTice	As piSST, but with monthly-varying SSTs and sea-ice taken from years 111-140 of each model's own abrupt4xCO2 experiment instead of from piControl.	Atmos-only	Year 111 of abrupt-4xCO2	30
a4SSTice-4xCO2	As a4SSTice, but CO2 is quadrupled, and the increase in CO2 is seen by both the radiation scheme and vegetation.	Atmos-only	Year 111 of abrupt-4xCO2	30
amip-a4SST-4xCO2	Same as amip, but a patterned SST anomaly is applied on top of the monthly-varying amip SSTs. This anomaly is a monthly climatology, taken from each model's own abrupt4xCO2 run minus piControl (using the mean of years 111-140 of abrupt4xCO2, and the parallel 30-year section of piControl). CO2 is quadrupled, and the increase in CO2 is seen by both the radiation scheme and vegetation.	Atmos-only	1979	36

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968 **Table 3.** Summary of CFMIP-OBS observational datasets available for comparison with COSP diagnostics.

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Dataset	Years	Observables	Applications	References
CALIPSO-GOCCP	2006/06 - 2012/10	Cloud fractions: 2D and 3D by phase. Scattering ratio histograms as function of height.	Vertical distributions of clouds. Cloud phase identification.	Chepfer et al., (2010); Cesana and Chepfer, (2013)
CloudSat	2006/06 - 2010/12	Reflectivity histograms as function of height.	Vertical distributions of clouds and precipitation	Marchand et al., (2009); Zhang et al., (2010)
ISCCP	1983/07-2008/06	Cloud top pressure – cloud optical depth histograms.	Cloud radiative properties. Long time series.	Rossow and Schiffer, (1999)
MODIS	2002/07 – 2015/11	Cloud top pressure – cloud optical depth histograms. Total, liquid and ice cloud fractions. Effective radius – optical depth histograms by cloud phase.	Cloud radiative properties. Effective size, and phase information.	Pincus et al., (2012); King et al., (2003)
MISR	2000/06 – 2013/05	Cloud top height (CTH) – cloud optical depth histograms	Cloud radiative properties. Independent estimate of cloud top height.	Marchand et al., (2010)
PARASOL	2003/05 - 2012/08	Monodirectional reflectance	Cloud radiative properties.	Konsta et al., (2015)

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## Main document changes and comments

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CFMIP is now entering its third phase, CFMIP-3, which will run in parallel with the current phase of the Coupled Model Intercomparison Project (CMIP6, Eyring et al., 2016)		
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the CFMIP-3/CMIP6 experiments and diagnostic outputs which constitute the CFMIP-3 contribution to CMIP6.		
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the CFMIP contribution to the current phase on the Coupled Model Intercomparison Project (CMIP6, Eyring et al., 2016).		
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and informal CFMIP-3 experiments which are organised independently of CMIP6. Please refer to the CFMIP website for announcements of these other initiatives and CFMIP annual meetings.		
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but for the purposes of this document CFMIP-3 should be considered to be synonymous with the CFMIP contribution to CMIP6.		
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(Given the names of other CMIP6 experiments this experiment might have been better named <i>amip-4xCO2-rad</i> , but this inconsistency was only noticed after the experiment names were finalised and propagated to the ESGF).		
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We also consider the time variation of feedbacks in *abrupt-4xCO2* experiments to be an important area to be investigated, as this can have a substantial impact on estimates of equilibrium sensitivity (e.g. Geoffroy et al., 2013). Andrews et al., 2015 investigated such effects using two atmosphere-only GCMs forced with SSTs and sea ice from their own *abrupt-4xCO2* experiments, and attributed the time variation in the feedbacks to changes in the pattern of surface warming. Pilot studies are ongoing to develop similar experiments based on a composite SST pattern response more representative of the CMIP5 ensemble mean. We plan to organise an informal pilot intercomparison based on this within CFMIP-3 and may subsequently propose these experiments as an extension to the CFMIP-3/CMIP6 experiment set.

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CFMIP-3

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CFMIP-3/CMIP6

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CFMIP-3

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CFMIP-3/CMIP6

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CFMIP-3

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CFMIP-3/CMIP6

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-3/CMIP6		
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-3/CMIP6		

-3/CMIP6

**Header and footer changes**

**Text Box changes**

**Header and footer text box changes**

**Footnote changes**

**Endnote changes**