Response to Reviewers Comment.

Reviewer 1 -

3 Mangeon et al. present the new fire and emissions schem ofr the Met Office's unified model. The approach presented has a reasonable complexity for to be useful in an Earth system model. The model is 4 evaluated using two different forcing datasets and different configurations of the ignition parameterization. Additionally the model performance is compared to the performance of fire weather 6 indices. Overall this is an interesting presentation suitable for publication within GMD. Nevertheless I have a number of suggestions which I believe will help to strengthen and improve the manuscript. 8 10 General comments: 11 The comparison with GFED focusses on stating that the emissions due to peat fires cannot be reproduced 12 by a model not including peatlands. This is correct, a solution could be to exclude the emissions from 13 14 peatlands from the comparison, as GFED provides the emissions for a number of different sources. 15 16 Reply: This is a fair point highlighted by the reviewer. Nevertheless, we do mention the figures from 17 GFED and have extended on that sentence. Changes to manuscript: 1342-345. Furthermore, 2002 and 2006 also saw important peat burning, with 18 GFEDv3 estimating peat emissions of 0.16 and 0.21 PgC respectively. In both of these years, the trend 19 in INFERNO differs from GFEDv3's (stagnation in 2002 and decrease in 2006). 20 21 22 I find the term fuel density to describe the amount of fuel per m-2 a bit confusing, as this term is often 23 used (for instance within spitfire) as the amount of fuel per volume. If it is the density per volume then the rate of spread decreases with increasing density. I would prefer the term fuel load. 24 25 26 Reply: We welcome the suggestion by the reviewer. Changes in manuscript: Across the manuscript, fuel density has been replaced with fuel load, and FD 27 28 29 30 A paragraph specifying the datasets used for the model evaluation is missing. 31 32 Reply: This paragraph was added to the manuscript at the end of the model configuration section. 33 Changes in manuscript: 1272-287. Evaluation was performed against the published data for GFEDv3, FINNv1, GFAS and GFEDv4. We also used the data from GFEDv4s (http://globalfiredata.org, 34 manuscript in preparation) and GFEDv4 (Giglio et al., 2013) to calculate grid-specific emissions and 35 burnt-area. The Global Fire Emissions Database (GFED) passes satellite observation of burnt area 36 37 through the Carnegie-Ames-Stanford-Approach (CASA) biogeochemical model in order to obtain

38 emissions from open burning. GFEDv4 (Giglio et al., 2013) innovates on GFEDv3 (Giglio et al., 2010) mainly through an updated algorithm to retrieve burnt area from MODIS satellite products and an 39 increased spatial and temporal resolution, to 0.25° and daily (this resolution was assessed in Mangeon et 40 al., 2015). Meanwhile GFEDv4s also includes the contribution from small fires (Randerson et al., 2012). 41 The Fire Inventory from NCAR version 1.0 (FINNv1, Wiedinmyer et al., 2011) provides high-resolution 42 (both temporal and spatial) global emissions of trace gas and particle emissions from open burning of 43 44 biomass. It focuses on rapid availability and assimilation in real time forecast and follows a similar 45 process to GFED to estimate emission, but its burnt area is obtained directly from fire pixel using land cover (Wiedinmyer et al., 2011). The Global Fire Assimilation System (GFAS, Kaiser et al., 2012), 46 unlike the aforementioned products, directly assess emissions from satellite-observed fire radiative 47 48 power more apt at detecting small fires and avoiding the uncertainty of biogeochemical models.

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The evaluation could also be a bit extended, for instance showing not only results of carbon emissions but also for the different chemical species.

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55 56 Reply: We have added total emissions for the different chemical species in Table 3. Still, we plan to further study specific species when performing the atmospheric composition evaluation of INFERNO, when coupled to the atmosphere. Further analysis will be included therein. Regarding the length of the evaluation, it has been increased following specific suggestions from the reviewers.

Changes to manuscript: 1373-381. Species-specific average emissions produced by the INFERNO 57 58 scheme are shown in Table 3 in Tg per year for the 1997-2010 period. CO and CH4 appear to be produced 59 in noticeably larger quantities than in observation-based emission estimates. This hints at an overrepresentation of smouldering-type combustion. In INFERNO this might be due to the emission 60 factors used, or the type of vegetation burnt. Table 3. Average annual emission (Tg per year) for 61 INFERNO with the interactive ignition mode and CRUNCEP reanalysis (3 - CRUNCEP) and the 62

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constant ignition mode and WFDEI reanalysis (1 - WFDEI), comparison to GFASv1 (Kaiser et al., 2012), GFEDv3 (van der Werf et al., 2010) and FINNv1 (Wiedinmyer et al., 2011) is provided.

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It remains unclear to me whether the fire model affects the vegetation dynamics, is there any tree mortality computed? also whether vegetation dynamics are included in the model simulations. If fire and vegetation dynamics interact a comparison of tree cover would be useful to evaluate that part of the model. If not, why don't they?

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Reply: At this stage, we focused on providing diagnostic tools with INFERNO. Having a vegetation that interacts with fires (carbon removal and tree mortality) might be investigated in the future. This would require a much deeper investigation and calibration of vegetation within JULES, beyond the scope of our work.

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Changes in manuscript: 1215-217. In its current early state, INFERNO provides a diagnostic tool, it does not remove carbon from vegetation nor does it lead to tree mortality. 1431-434. To close the vegetation-

77 fire feedback, INFERNO will eventually need to remove carbon from vegetation and to include tree 78 mortality. 79 Specific comments: 80 81 82 1. 19: you could add the outcome of the fire index diagnostics comparison. 83 84 Reply: Thanks Changes to manuscript: 119-20. We found INFERNO captured global burnt area variability better than 85 86 individual indices, and these performed best for their native regions. 87 88 1. 21: is this spatial or temporal correlation? Is it significant? 89 Reply: This is spatial correlation. We have modified the latter text (to keep the abstract short). 90 Changes to manuscript: 1293-294. we found a spatial correlation of R=0.66 for burnt area and R=0.59 91 92 for emitted carbon, both passing the t-test with 95% significance. 93 94 1. 101: the scaling factor is the 7.7, please specify. 95 Reply: This is now specified in the text. 96 Changes in manuscript: 1. 102. Equation 2 includes a scaling factor of 7.7 ([...] 97 98 99 1. 102-5: if you assume fNS=1, you don't need it in the equation, adding this assumption after presenting 100 the equation might be more clear: total ignitions can be represented as: eq3, here fNS equals 1 for mode 101 1 and 3 and follows eq. 2 for mode 3. 102 103 Reply: Thanks for the comment, we clarified it as suggested in the manuscript. 104 Changes to the manuscript: 1103-107. Total ignitions [...] Here fNS = 1 for mode 1 and 2 and follows 105 eq. 2 for mode 3. 106 107 1. 117: Leaf carbon is the living biomass? 108 109 Reply: Yes - this was clarified as aboveground to only represent living leaves Changes in the manuscript: 1119. (LeafC,PFT, aboveground) 110 111 112 1. 119: I think this should say FDPFT, the equation actually does not scale lienarly between 0 and 1, it

jumps from 0 to $Fuel_{low}/(Fuel_{high}$ - $Fuel_{low}$). I guess the equation should be

 $((FPMc + leafC) - Fuel_{low}) / (Fuel_{high} - Fuel_{low}). \ Additionally, \ the \ equation \ is \ not \ defined \ for \ fuel \ density \ densi$

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being equal to fuellow and fuelhigh.

- 117 Reply: A great thanks to the reviewer for spotting this, the equation was changed to reflect this.
- 118 Changes in the manuscript: 1118. Eq. 5.

$$FL_{PFT} = \begin{cases} 1 \text{ for } Fuel_{high} < (DPM_C + Leaf_{C,PFT}) \\ \frac{(DPM_C + Leaf_{C,PFT}) - Fuel_{low}}{Fuel_{high} - Fuel_{low}} \text{ for } Fuel_{low} \leq (DPM_C + Leaf_{C,PFT}) \leq Fuel_{high} \\ 0 \text{ for } Fuel_{low} > (DPM_C + Leaf_{C,PFT}) \end{cases}$$

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- 121 l. 123: Here again, the normalization term should be $(RH RH_{low})/(RH_{up} RH_{low})$, C2 please rewrite
- 122 the equation similar to eq.5 to define the bounds for relative humidity being higher and lower then the
- 123 thresholds

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- 125 Reply: Again, thanks to the reviewer for spotting this and suggesting the edit
- 126 Changes in the manuscript: 1125.

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$$F_{PFT} = \begin{cases} e^* \frac{RH - RH_{low}}{RH_{up} - RH_{low}} e^{-2R} FL_{PFT} (1 - \theta) & \text{for } RH_{up} < RH \\ e^* \frac{e^* e^{-2R} FL_{PFT} (1 - \theta)}{RH_{low}} & \text{for } RH_{low} \leq RH \leq RH_{high} \\ 0 & \text{for } RH_{low} > RH \end{cases}$$

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129 1. 127: FD is Fuel Density or Fuel density index? 1. 126-127: which formula are you referring to?

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- 131 Reply: FD (or now FL) stands for fuel load index here (this was clarified). We also clarified the term
- 132 further to avoid confusion.
- 133 Changes to the manuscript: 1129. fuel load (DPMc and Leafc,pft). (l. 153) Fuel Loaf index (FL)

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- 135 1. 133: explain how the average burnt area was determined. There is no difference between temperate
- and tropical trees?

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- 138 Reply: We specified the average burnt area was heuristically determined. Far from stating this method
- is perfect, we feel within this paragraph and the next we give enough justification for it while also
- 140 presenting the alternatives. Temperate and tropical trees are assigned the same average burnt area, for
- 141 simplicity
- 142 Changes to the manuscript: 1137. Sub-categories of trees, grass and shrubs are not differentiated

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- 144 l. 146: That suggests you should vary your pft specific burned area. Are there any indications in your
- results that this is necessary? it might be rather a point for the discussion of your results.

- 147 Reply: This is difficult to ascertain without comparing earlier components (fire count or frequency).
- 148 Which we have not done for this paper (and INFERNO does not have the capability to produce fire count
- estimates). Thank you to the reviewer for suggesting the discussion was a more appropriate location for
- this point, which we have also extended upon.
- 151 Changes to manuscript: 1307-311. INFERNO's approach to burnt area only considers trees, grass and
- shrub cover and was determined heuristically, meanwhile Hantson et al. (2014) found global fire size

- was mostly influenced by precipitation, aridity and human activity (population density and croplands).

 Further parameterizations for fire size exist (e.g., Hantson et al., 2015, 2016) which could improve
 INFERNO burnt area estimates while maintaining simplicity and traceability.
- 157 1. 154: CCmin and CCmax are the same for leaves and stems.

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159 Reply: Again, thanks to the reviewer for spotting this. CC (stem) was modified accordingly.

Changes to the manuscript: 1157. [...] and stems (CCmin = 0.0 and CCmax = 0.4) [...]

- 162 1. 158: I don't see why this is makes it justifyable. more interesting would be why you changed the value,
- 163 was it to tune the emissions?164
- Reply: We realize stating 'justifiable' implies GFED's assumptions should be mimicked. We have rephrased it to convey the message more effectively. We also go into more depth on the matter.
- 167 <u>Changes to the manuscript:</u> 1160-164. Nevertheless, GFED uses a more complex representation of 168 moisture across multiple fuel types and only accounts for fires that were observed. In comparison, our 169 scheme only relies on soil moisture and was much more sensitive to minimum combustion, such that the
- 170 contribution from moist forested areas (e.g., rainforests) needed to be reduced by increasing the impact 171 of soil moisture (reducing stems' CC_{min}).
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 1. 200: what happens with population and lightning flash rates if JULES is not used in standalone version.
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 175 Reply: Ignition mode 2 is essentially the coupled version (to be refined and later submitted for publication): the lightning can be interactively simulated (population still needs to be prescribed or assumed as constant).
- 178 <u>Changes to the manuscript:</u> 1211-212. Interestingly, lightning can be interactively simulated in
- atmospheric models (not population), although this will not be explored in this paper.
- 181 1. 206-225: Give equations for the fire weather indices.
- 182 Reply: We have added equations for the more straightforward McArthur and Nesterov index. For the
- 183 Canadian FWI we have only included the key equation, the full index is very complex and we refer back
- to the original papers for a detailed description.
- 185 <u>Changes to manuscript:</u> 1226-250. Equation 12-15 and related details in section 2.3. Appendix A now
- 186 contains the equations used in the Canadian Fire Weather Index.187
- 1. 254: How is the correlation computed? spatial or temporal? if temporal, is the correlation computed
 for each grid cell or just for the global total? pleave give significance levels.

Reply: Similar to previous point. This is spatial correlation, which is computed for each gridcell. We found p-values of 0 for both correlations (and these passed the t-test with 95% significance level). There are about 28000 grid-cells included in the analysis which explains the virtually 0 p-value here.

Changes to manuscript: 1292-296. INFERNO accurately diagnoses total fire occurrence and emissions over the 1997-2010 period: we found a spatial correlation of R=0.66 for burnt area and R=0.59 for emitted carbon, both passing the t-test with 95% significance.

 I. 258-262: I don't think the gridcell with maximum burned area is an important bench-mark. But what about seasonality? Emissions for the different sources given by GFED could also be interesting. or burned area separated for grass and woody pfts.

Reply: We wrote the maximum burnt area for two reasons: 1. It mirrors the emitted carbon, where the maximum is much more indicative of a model bias (with peat). 2. It remains important to assess whether INFERNO can assess extreme burnt area, and it adds a narrative to the article. Therefore, we would prefer to keep this point, while expanding others as suggested. Nevertheless, we will follow the suggestions of the reviewer and mention the main PFTs that contribute to burnt area, and emissions (C4 grass and Broadleaf Evergreen Tree (Tropical)). More details were also added for peat-specific GFED emissions (see previous comment response).

208 emissions (see previous comment response).
209 Changes to manuscript; 1299-301. Over the s

<u>Changes to manuscript:</u> 1299-301. Over the studied period, C4 grass were the main contributors to burnt area in INFERNO (a mean 2.34 Mkm² per year), meanwhile Broadleaf Evergreen Trees (Tropical) led to the most emitted carbon (a mean 1.48 Pg per year).

l. 264: the peat emissions given by GFED could be excluded here, are crop fires and emissions due to deforestation actually somehow included in the model? otherwise they could also be excluded.

Reply: Although these are not explicitly represented in the model, the function that estimates ignitions according to population density represents the various ways humans can interact with fires – including deforestation and crop fires (although in these simulations, crops are equivalent to grass). Peat and crop/deforestation fire emissions are different in that the latter are somehow accounted for (albeit not explicitly), while the former simply is not present in the land surface model, and thus in INFERNO. We detail the contribution of peat fires to GFED later in the discussion, and have expanded on that analysis. Changes to manuscript: 1342-345. Furthermore, 2002 and 2006 also saw important peat burning, with GFEDv3 estimating peat emissions of 0.16 and 0.21 PgC respectively. In both of these years, the trend in INFERNO differs from GFEDv3's (stagnation in 2002 and decrease in 2006).

1. 338: interestingly the mid latitudes are not well captured by the fire weather indices, might be the human influence? Including the other ignition modes of the model could give an indication why the model in better than the indices. Any significance values on the correlation?

230 Reply: Thanks the reviewer for this comment – we performed further analysis with ignition mode 3 and found a correlation coefficient of R=0.1617 (lower than mode 1), suggesting our scheme for interactive 231 232 human ignition is not able to improve estimates at mid-latitudes. We also performed a t-test (95% significance) on the values for correlations, and they were all passed successfully but for global and high 233 latitudes for the McArthur index, and mid-latitude for the Nesterov index. 234 Changes to manuscript: Italics mean the correlation was not significant (p-value above 0.05). 235 236 237 1. 361: where did you show that the precipitation has an important impact? 238 239 Reply: During the analysis of our results we found precipitation varied significantly between the two driving datasets (CRUNCEP and WFDEI-GPCC). Still, we do not wish to do a full analysis of the causes 240 241 for this discrepancy and its impact on INFERNO here, and chose to simply remove this confusing term (indeed Fig. 3 and Table 2 both show the impact of using a different meteorological dataset). 242 243 Changes to manuscript: 1418. The use of different present-day meteorological datasets has an important 244 impact on the magnitude and variability of our diagnostics. 245 246 1.370: You assessed the uncertainty of the ignitions by including the different ignition modes, but how does this dampen the impact of this uncertainty in inferno? 247 248 Reply: Thank you for the comment – the language was poorly chosen, and the purpose of these multiple 249 250 ignition is to assess rather than dampen indeed. 251 Changes to manuscript: 1428-429. Accordingly, we include different modes to examine the impact of 252 ignitions (human or natural) in INFERNO. 253 254 1. 375: what do you mean by vaporized? 255 256 Reply: Thanks to the reviewer for this comment - vaporized was not the most accurate term here, we 257 modified the sentence accordingly. 258 Changes to manuscript: 1434-436. For instance, litter influences flammability but only live vegetation 259 leads to emissions while in reality litter significantly contributes to observed fuel consumption (van 260 Leeuwen et al., 2014). 261 262 1. 379-382: I don't understand. what do you mean by INFERNO's meteorological and hydrological 263 assimilation? In what sense are the other fire schemes more specialized? 264 Reply: This was referring to the way INFERNO assimilates weather and water. However, the goal of 265 this sentence is to advise anyone that would develop INFERNO further to keep its 'simplicity and 266 267 linearity'. We have changed the sentence to reflect this. What we meant by specialized fire scheme is more often referred to as 'process-based models'. 268

- 269 <u>Changes to manuscript:</u> 1440-441. The value of this model being its simplicity and linearity, any improvements to INFERNO should follow this vision; complex parameterizations are better suited for
- process-based fire schemes (e.g., Lasslop et al., 2014; Li et al., 2013, p.1).

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Fig.A1: label the subpanels. why does the temperature function not scale between 0 and 1?

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- 275 Regarding the temperature, we chose to restrict the display here to the range of realistic
- 276 temperatures observed on Earth (similarly, precipitation and fuel density are restricted to the 'key'
- location of their representative functions). Regarding the request to label the subpanels it's now done.
- 278 Changes to manuscript: 1452. Updated Fig. A1 and changed its caption to:
- 279 Fig. A1. The mathematical functions used for individual dependencies of INFERNO on key driving
- variables for flammability (a,b,c,d,e) and ignitions (f), within the range of reasonable earth observations.
- 281 Note the population density only influences the model output if ignition mode 3 is selected (interactive
- 282 lightning and human ignition).

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- Reviewer 2 -
- 285 Review
- 286 INFERNO: A fire and emissions scheme for the Met Office's Unified Model
- Mangeon et al. This paper describes a simplified model that projects biomass burning activity, burnt
- area, and emissions globally. The model framework uses climatic and meteorological inputs and land
- 289 cover characteristics to drive the emissions model. The ignition sources can be varied, and, for the
- 290 purposes of the model evaluation presented, are prescribed three different ways in order to assess the
- 291 sensitivity of the model to different ignition parameterizations. The land cover inputs are provided by
- 292 the JULES model. The fire model is run for current conditions and compared with other fire model
- 293 outputs (primarily GFED).

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- 295 This paper is written extremely well, and the modeling tool described is a unique contribution. It uses
- different approaches than other available models that project fires in global climate models and will be
- 297 a useful tool to be incorporated within the UK Met Office's Earth System Model. The assumptions made
- 298 in the model parameterizations are reasonable and well justified throughout. The manuscript is very
- 299 appropriate for Geoscientific Model Development, and I recommend publication after only minor
- 300 comments that I provide here.

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- 302 General Comments:
- 303 Section 2.2: How are the PFTs allocated within each grid cell? This is not explained thoroughly in Section
- 304 2.2, and the paper cited as a reference is still "in prep".

306 Reply: Given the paper is about fire and there's no feedback on vegetation, we would prefer not to detail 307 this further. However, we have made this more explicit in the manuscript and the reference (now accepted 308 Changes in manuscript: 1293-195. Fractional coverage of PFTs in any gridcell is based on competition 309 for resources (light and water), and governed by Lotka-Volterra competition equations, and based on a 310 311 tree-shrub-grass dominance hierarchy (Cox, 2001). 312 313 Section 2.3: Emissions from Akagi et al. (2011) have been updated and can be incorporated within future versions (see Section 3 at http://bai.acom.ucar.edu/Data/fire/). 314 315 316 Reply: Now mentioned in the text 317 Change in the manuscript: 1168-169. Section 2.1.5, added: "Updated EFs for Akagi et al. (2011) were not used in this version of INFERNO, these can be found in section 3 of: 318 319 http://bai.acom.ucar.edu/Data/fire/" 320 321 Section 4: I was confused about the fact that there were two different versions of "GFEDv4" used to 322 evaluate the INFERNO estimates. Can this be made more clearly? (i.e., better define and label the two different outputs)? 323 324 325 Reply: Clarification added in new paragraph on datasets used for comparison (which describes datasets 326 compared to). Also improved the legend in figure 3 (v4 + small fires). Changes in manuscript: 1272-287 (new paragraph) and Fig. 3 GFEDv4s (v4 + small fires). 327 328 329 Lines 255 and following sentences: Please clarify which model (INFERNO v GFED) was higher/lower. For example, Line 255 can be re-written: 330 331 "We notice that the burnt area predicted by INFERNO is higher in all regions other than Australia and New Zealand, and southern hemisphere Africa when compared to 332 GFED4." 333 334 335 Reply: Rephrased this sentence to clarify the comparison 336 Changes in manuscript: 1295-296. Compared to GFEDv4, we notice INFERNO estimates higher burnt area in all regions apart from Australia and New Zealand, and southern hemisphere Africa. 337 338 339 Paragraph starting at line 325: Is it possible to compare the fire indices calculated here with the real data 340 for current conditions?

Reply: Fire indices were developed as a means to assess fire dangers, within specific biomes, while they

did not aim to assess 'real data' like burnt area. However, by assessing the R-coefficient with GFED burnt area we study the variability of the indices, and their capacity to mimic the variability observed in

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'real data'. This comparison makes fire indices analogous to INFERNO's flammability, which in ignition mode 1 (used in this comparison with fire indices), is the only source of variability.
 Changes in manuscript: 1386-388. This analysis relies on the assumption that fire indices can be used as a proxy for the variability of fire occurrence and spread, and eventually of burnt area (not the magnitude).
 Figure 3: Why are not GFAS and FINN outputs compared in both panels of the figure?

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- 352 Reply: Neither of the original papers presented results for the other category, i.e. FINN
- 353 (http://www.geosci-model-dev.net/4/625/2011/gmd-4-625-2011.html) does not contain information on
- 354 Emitted Carbon (although it does on Biomass burned, conversion is not obvious). Meanwhile GFAS
- does not estimate burnt area. Accordingly, we also use GFAS, GFEDv3 and FINN to examine specific
- 356 species
- 357 <u>Changes to the manuscript:</u> Table 3. & 1272. Evaluation was performed against the published data for
- 358 GFEDv3, FINNv1, GFAS and GFEDv4.

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- 360 Editorial Comments:
- 361 Title: The UK Met Office should be defined in the title.
- 362 Reply: The title now includes UK. However due to branding Met Office should remain so (rather than
- 363 Meteorological Office).
- 364 Changes to manuscript: title: 11. INFERNO: a fire and emissions scheme for the UK Met Office's Unified
- 365 Model.

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- Lines 59 and 60: The present tense should be applied (i.e., change "used" to "uses")
- 368 Reply: OK
- 369 Changes to manuscript: 160-62. In short, that parameterization uses monthly mean temperature, relative
- 370 humidity and precipitation to simulate fuel flammability. It also uses human population density and
- 371 lightning to represent ignitions.

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- 373 Line 67: Add a comma after (Ex)
- 374 Reply: OK
- 375 Changes to manuscript: 168 (Ex), and

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Line 168: It may be useful to let the reader know that [C] will be described in the next section.

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- 379 Reply: To an extent we felt the next section did not describe this well enough, therefore we modified this
- 380 sentence to be more explicit.
- Changes to manuscript: 1175-176. and [C] is the dry biomass carbon content which we assume as 50%
- 382 (a common simplification; Lamlom and Savidge, 2003).

384 Line 180 (and elsewhere): When "which" is used, there should be a comma preceding it. In this case, 385 there should be a comma after "(see Eq. 6.8)" 386 Reply: Yes, because the clause is restrictive here, "which" should be preceded by a comma. 387 388 Changes to manuscript: 1187. (see Eq. 6.8), which 389 Line 182: Layers should be plural 390 391 392 Reply: Yes, thanks 393 Changes to manuscript: 1189. Layers 394 395 Line 258: I recommend changing "observes" to "projects" (or something like that). 396 397 Reply: Thanks to the reviewer for this suggestion that we applied. Changes to manuscript: 1301. "GFEDv4 projects the" ... note: we have also changed all mentions of 398 399 GFED4 to GFEDv4 for consistency. 400 401 Line 328: condition should be plural 402 403 Reply: Yes 404 Changes to manuscript: 1385. same vegetation conditions 405 406 Line 365-6: Should this be "...presence of an anthropogenic ignition source." 407 408 Reply: The point we were trying to make would apply to both anthropogenic and lightning-started fires. 409 Albeit it was not clear, which we tried to improve (see changes below). Changes to manuscript: 1422-424. Similarly, lightning appears to more frequently ignite fires in wet 410 411 environments (rainforests) while flammable environments (savannah, grasslands) with rarer lightning 412 are sensitive to the presence of an anthropogenic ignition source. 413 414 Line 375-376: a citation should be given. 415 416 Reply: The sentence was slightly changed upon reading the study now cited (van Leeuwen et al., 2014) 417 and to improve its flow. Changes to manuscript: 1435. The sentence now reads: "For instance, litter influences flammability but 418 419 only live vegetation is vaporized while in reality litter significantly contributes to observed fuel

consumption (van Leeuwen et al., 2014)."

INFERNO: a fire and emissions scheme for the **UK** Met

2 Office's Unified Model

- 3 Stephane Mangeon^{1,2}, Apostolos Voulgarakis¹, Richard Gilham², Anna Harper³,
- 4 Stephen Sitch⁴, Gerd Folberth²
- Department of Physics, Imperial College London, London, United Kingdom
- 6 ²Met Office, FitzRoy Road, Exeter, EX1 3PB, UK
- 7 College of Engineering, Mathematics, and Physical Sciences, University of Exeter, Exeter, UK
- 8 College of Life and Environmental Sciences, University of Exeter, Exeter, UK
- 9 Correspondence to: Stéphane Mangeon (stephane.mangeon12@imperial.ac.uk)
- 10 Abstract. Warm and dry climatological conditions favour the occurrence of forest fires. These fires then
- become a significant emission source to the atmosphere. Despite this global importance, fires are a local
- 12 phenomenon and are difficult to represent in a large-scale Earth System Model (ESM). To address this,
- $13 \qquad \text{the INteractive Fire and Emission algoRithm for Natural envirOnments (INFERNO) was developed.} \\$
- 14 INFERNO follows a reduced complexity approach and is intended for decadal to centennial scale climate
- 15 simulations and assessment models for policy making. Fuel flammability is simulated using temperature,
- relative humidity, fuel <u>load</u> as well as precipitation and soil moisture. Combining flammability with
- 17 ignitions and vegetation, burnt area is diagnosed. Emissions of carbon and key species are estimated
- 18 using the carbon scheme in the JULES land surface model. JULES also possesses fire index diagnostics
- 19 which we document and compare with our fire scheme. We found INFERNO captured global burnt area
- 20 variability better than individual indices, and these performed best for their native regions. Two
- 21 meteorology datasets and three ignition modes are used to validate the model. INFERNO is shown to
- increorology datasets and three ignition modes are used to various the model. In ERNO is shown to
- 22 effectively diagnose global fire occurrence (R=0.66) and emissions (R=0.59) through an approach
- 23 appropriate to the complexity of an ESM, although regional biases remain.

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Deleted: density

26 1 Introduction

- 27 Fire is a key interaction between the atmosphere and the land surface (Bowman et al., 2009). Its impacts
- are wide-ranging: it influences forest succession (Bond and Keeley, 2005), is a tool for deforestation
- 29 (van der Werf et al., 2009) and is an important natural carbon source (Bowman et al., 2013), while it also
- 30 provides a major natural hazard to humans through property and infrastructure destruction and air quality
- degradation (Johnston et al., 2012; Marlier et al., 2013). Not only are biomass burning emissions
- 32 substantial in magnitude (Lamarque et al., 2010), they also drive the variability of atmospheric
- 33 composition (Spracklen et al., 2007; Voulgarakis et al., 2010, 2015) and impact short-term climate
- 34 forcing (Tosca et al., 2013).
- 35 There are feedbacks between fire and climate: low-humidity conditions cause droughts, which enhance
- 36 fire activity (Field et al., 2009), which, in turn, emits aerosols and trace gases (Akagi et al., 2011),
- 37 influencing the abundances of radiatively active atmospheric constituents, cloud formation and lifetime,
- and in turn precipitation, and surface albedo (Voulgarakis and Field, 2015). Bistinas et al. (2014) showed
- 39 global fire frequency is correlated with land-use, vegetation type and meteorological factors (dry days,
- 40 soil moisture and maximum temperature) and human presence tends to noticeably reduce fire activity
- 41 (land-management, landscape fragmentation and urbanization). Examining and quantifying such impacts
- 42 and feedbacks is paramount to Earth System Models (ESMs), yet to integrate vegetation fires presents
- 43 many challenges as it intricately links multiple disciplines from ecology to atmospheric chemistry and
- 44 physics and climate science.
- 45 Integration of fires into Dynamic Global Vegetation Models (DGVMs) was the first step towards fire
- 46 within ESMs (e.g. (Arora and Boer, 2005; Fosberg et al., 1999; Li et al., 2012; Pfeiffer et al., 2013; Sitch
- 47 et al., 2003; Thonicke et al., 2001, 2010; Venevsky et al., 2002; Yue et al., 2014). Vegetation fires have
- 48 been implemented into only a few ESMs, e.g. ECHAM (Lasslop et al., 2014) and the Community ESM
- 49 (Li et al., 2013, 2014, p.2).
- 50 Here, we present and evaluate the INteractive Fire and Emission algoRithm for Natural envirOnments
- 51 (INFERNO) and its implementation. INFERNO is a necessarily simple parameterization that focuses on
- 52 the large-scale occurrence of fires and is suitable for ESM application. The model uses a few key driving
- variables while retaining a broadly accurate parameterization for fire emissions. INFERNO's
- 54 performance against observations and well established and operationally relevant fire indices is
- 55 presented.

56 2 Model description

57 **2.1 INFERNO**

- 58 INFERNO was constructed upon the simplified parameterization for fire counts proposed and evaluated
- 59 for the present-day by (Pechony and Shindell, 2009), which was subsequently shown to provide a good
- 60 estimate for large-scale fire variability over climatological timescales (Pechony and Shindell, 2010). In
- short, that parameterization uses monthly mean temperature, relative humidity and precipitation to
- 62 simulate fuel flammability. It also uses human population density and lightning to represent ignitions.
- 63 To incorporate this parameterization within the Joint UK Land Environment Simulator (JULES, Best et

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al., 2011; Clark et al., 2011), several changes were applied. Upper layer soil moisture is used to represent precipitation memory while precipitation acts as a rapid fire deterrent. Vegetation Density was replaced by Fuel Load index, dependent on leaf carbon and Decomposable Plant Material (DPM), i.e. litter. Such a relationship with fine fuel and moisture was used in Thonicke et al. (2001). Furthermore, we developed a parameterization to obtain burnt area (BA), emitted carbon (EC) and fire emissions of different species $(E_X)_2$ and our fire diagnostics are made for each of the nine Plant Functional Types (PFTs) in the current version of JULES (Harper et al., 2016).

Figure 1 summarizes the mechanisms of INFERNO, and Fig. A1 illustrates the dependence of INFERNO on individual driving variables.

Ignitions

Lightning Strikes

Population Density (Standalone only)

Fuel Density

Precipitation and Soil Moisture

Fire Emissions

Fire Emissions

Fire Emissions

Fig. 1. Schematic summarizing the INteractive Fire and Emission algoRithm for Natural envirOnments (INFERNO) and its key components and behaviour. Ignitions can be accounted for in a variety of ways (see Sect. 2.1.1), meteorology influences flammability (see Sect. 2.1.2), while plant coverage influences burnt area (see Sect. 2.1.3), finally emissions are calculated according to leaf and stem carbon for each PFT (see Sect. 2.1.4).

2.1.1 Ignitions (*I*)

82 INFERNO calculates ignitions in either one of three modes:

First, we can assume constant or ubiquitous ignitions, currently calibrated to a global average of $I_T = 1.67$ ignitions km⁻² month⁻¹. This corresponds to 1.5 ignitions km⁻² month⁻¹ due to humans (I_A), heuristically determined, and 0.17 ignitions km⁻² month⁻¹ natural ignitions due to lightning (I_N), derived from the multi-year annual mean of 2.7 strikes km⁻² year⁻¹ (Huntrieser et al., 2007) assuming 75% of strikes being cloud-to-ground (Prentice and Mackerras, 1977). This mode inherently suppresses the variability in fires due to any anthropogenic or natural ignition changes (Pechony and Shindell, 2009, 2010).

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- 93 Second, human ignitions and suppressions can be assumed to remain constant at the global mean value
- mentioned above ($I_A = 1.5 \text{ ignitions km}^{-2} \text{ month}^{-1}$), however cloud-to-ground lightning strikes may vary,
- 95 and in addition each strike is assumed to start a fire. This mode accounts for natural variability in fire
- 96 ignitions, which can be simulated within an ESM, or prescribed from observations.
- 97 Third, varying human ignitions and suppressions and varying natural ignitions (cloud-to-ground
- 98 lightning strikes, as in mode 2). This was the original ignition approach in Pechony and Shindell (2009),
- 99 which was left unchanged and is detailed below. In this ignition mode, anthropogenic ignition and
- suppression depends on population density (PD), as proposed by Venevsky et al. (2002).

$$I_A = k(PD) PD \alpha \tag{1}$$

- 102 PD is in units of people km⁻², and $k(PD) = 6.8 \times PD^{-0.6}$ is a function that represents the varying
- 103 anthropogenic influence on ignitions in rural versus urban environments. The parameter $\alpha = 0.03$
- 104 represents the number of potential ignition sources per person per month per km². Both natural and
- anthropogenic ignitions have the potential to be suppressed by humans, such that the fraction of fires not
- 106 suppressed is:

107
$$f_{NS} = 7.7 (0.05 + 0.9 \times e^{-0.05 PD})$$
 (2

- Equation 2 includes a scaling factor of 7.7 (Pechony and Shindell, 2009) originally introduced to calibrate
- the number of fires to MODIS observations. Total ignitions (I_T , in units, ignitions m⁻² s⁻¹) can be
- 110 represented as (Eq. 3):

111
$$I_T = (I_N + I_A) f_{NS} / (8.64 \times 10^{10})$$
 (3)

- Here $f_{NS} = 1$ for mode 1 and 2, and follows eq. 2 for mode 3. Dividing by 8.64×10^{10} converts
- 113 ignitions km⁻² month⁻¹ to ignitions m⁻² s⁻¹.

114 **2.1.2 Flammability** (*F*)

- 115 We adapt the (Pechony and Shindell, 2009) scheme for flammability to function interactively within an
- 116 ESM (see Eq. 6). Starting from the saturation vapour pressure (e*, Eq. 4; Goff and Gratch, 1946) and
- 117 its temperature dependence, we introduce a Fuel Load index (FL. F.T., Eq. 5) as well as Relative Humidity
- 118 (RH), precipitation and soil moisture in order to obtain Flammability (Eq. 6). The land surface model
- [119] (JULES) determines soil moisture content (θ) and fuel $load (DPM_{C_a} and Leaf_{C,PFT})$

$$\log_{10}\left(e^{*}\right) = a\left(\frac{T_{s}}{T} - 1\right) + b\log_{10}\left(\frac{T_{s}}{T}\right) + c\left(10^{d\left(1 - \frac{T_{s}}{T}\right)} - 1\right) + f\left(10^{h\left(\frac{T_{s}}{T} - 1\right)} - 1\right) \tag{4}$$

- As illustrated in Eq. 4, INFERNO utilizes temperature (T in K, at 1.5 m height). The Goff-Gratch (Eq.
- 4) uses the constants: a = -7.90298; b = 5.02808; $c = -1.3816 * 10^{-7}$; d = 11.344; $f = 8.1328 * 10^{-7}$
- 123 10^{-3} ; h = -3.49149 and the water boiling point temperature $T_s = 373.16$ K.

124
$$FL_{PFT} = \begin{cases} 1 \text{ for } Fuel_{high} < (DPM_C + Leaf_{C,PFT}) \\ \frac{(DPM_C + Leaf_{C,PFT}) - Fuel_{low}}{Fuel_{high} - Fuel_{low}} \text{ for } Fuel_{low} \le (DPM_C + Leaf_{C,PFT}) \le Fuel_{high} \\ 0 \text{ for } Fuel_{low} > (DPM_C + Leaf_{C,PFT}) \end{cases}$$

$$(5)$$

- 125 Equation 5 shows FL_{PFT} is taken as the PFT-specific leaf carbon (Leaf_{C,PFT}, aboveground) plus the
- 126 carbon within decomposable plant material (DPM_C) . DPM is a soil carbon pool of which we assume
- 70% is available to fires i.e. near-surface (DPM is shared across all PFTs). FL scales linearly between 0
- (at a threshold of $Fuel_{low} = 0.02 \text{ kgC m}^{-2}$) and 1 (at a threshold of $Fuel_{high} = 0.2 \text{ kgC m}^{-2}$). Similar

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147 approaches to represent fuel availability within fire parameterizations have commonly been adopted

148 (Arora and Boer, 2005; Li et al., 2012; Thonicke et al., 2010).

$$F_{PFT} = \begin{cases} e^* \frac{e^* e^{-2R} F L_{PFT} \left(1 - \theta\right) \text{ for } R H_{up} < RH}{R H_{up} - R H_{low}} e^{-2R} F L_{PFT} \left(1 - \theta\right) & \text{for } R H_{low} \le RH \le R H_{high} \\ 0 \text{ for } R H_{low} > RH \end{cases}$$

150 RH is the relative humidity (%) and R is the precipitation rate (mm day-1). The influence of relative

humidity (RH) scales between (and is bound by): 0 (at a threshold of $RH_{low} = 10\%$) and 1 (at a threshold 151

152 of $RH_{up} = 90\%$). We then adapt the formula by replacing a vegetation index dependent on leaf area

153 index with the Fuel Load index (FL). Finally, Flammability (F_{PFT}) is dependent on upper-level (down

154 to 0.1 m) soil moisture: θ is the unfrozen soil moisture as a fraction of saturation. The individual

155 importance of these variables to our model is illustrated in Fig. A1.

2.1.3 Burnt Area (BA)

Our approach is to associate an average burnt area per fire to each PFT, effectively decoupling the fire-157

spread stage from local meteorology and topography, which is typically not resolved in the relatively

coarse grid of an ESM. An average burnt area $(\overline{BA_{PFT}})$ was heuristically determined for each PFT: 0.6, 159

160 1.4 and 1.2 km² for trees, grass and shrubs, respectively, such that grass and shrubs will fuel larger fires

161 than trees. Sub-categories of trees, grass and shrubs are not differentiated. Observational evidence

162 supports that the land cover type is an efficient way to characterize fires, which tend to be larger in

163 grasslands than in forests (Chuvieco et al., 2008; Giglio et al., 2013). The BA is then calculated following

164 Eq. 7:

156

158

$$165 BA_{PFT} = I_T F_{PFT} \overline{BA_{PFT}} (7)$$

Here BA_{PFT} is the burnt area (fraction of PFT cover burnt per second) for each PFT; meanwhile the 166

167 number of ignitions times the flammability $(I_T F_{PFT})$ represents the number of fires.

168 Inferring burnt area from number of fires in this manner stands out from other fire models which utilize

169 wind speed (Arora and Boer, 2005; Thonicke et al., 2010; Li et al., 2012), effectively modelling the fire

170 rate of spread. Wind is key to the modelling of individual fires; yet implementing wind effectively within

171 fire models designed for the relatively coarse grid of ESMs was found to be problematic (Lasslop et al.,

172 2014, 2015). Conversely, Hantson et al. (2014) found global fire size was mostly influenced by

173 precipitation, aridity and human activity (population density and croplands).

174 2.1.4 Emitted Carbon (EC)

175 To account for the wetness of fuel in INFERNO, combustion completeness (the fraction of biomass

176 exposed to a fire that was volatized) scales linearly with soil moisture (as a fraction of saturation) with

different upper and lower boundaries for leaf and stem carbon. 177

178
$$EC_{PFT} = BA_{PFT} \sum_{leaf,stem}^{i} \left(CC_{min,i} + \left(CC_{max,i} - CC_{min,i} \right) (1 - \theta) \right) C_i$$
 (8)

Equation 8 shows how the PFT-specific emitted carbon (EC, in kgC m⁻² s⁻¹) is computed. BA is the burnt 179

180 area (fraction s⁻¹), CC_{min} and CC_{max} are the minimum and maximum combustion completeness for both

181 leaves ($CC_{min} = 0.8$ and $CC_{max} = 1.0$) and stems ($CC_{min} = 0.0$ and $CC_{max} = 0.4$), C_i is the carbon

stored in each PFT's leaves or stems (kgC m⁻²). The parameters used for combustion completeness 182

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194 $(CC_{min}$ and $CC_{max})$ are similar to the Global Fire Emission Database (GFED; yan der Werf et al., 2010),

albeit with lower minimum combustion of stems (0.0 as opposed to 0.2). Nevertheless, GFED uses a

196 more complex representation of moisture across multiple fuel types and only accounts for fires that were

197 observed. In comparison, our scheme only relies on soil moisture and was much more sensitive to

198 minimum combustion, such that the contribution from moist forested areas (e.g., rainforests) needed to

199 be reduced by increasing the impact of soil moisture (reducing stems' CCC_{min}).

2.1.5 Emitted Species (E_X)

195

200

202

201 There has been a significant amount of work on estimating emission factors (EFs) across fire biomes

(such as savannahs, boreal forest etc.). This was synthesized in Akagi et al. (2011) as well as Andreae

203 and Merlet (2001) and its updates. Updated EFs for Akagi et al. (2011) were not used in this version of

204 INFERNO, these can be found in section 3 of: http://bai.acom.ucar.edu/Data/fire/. To convert biome-

specific EFs to PFT specific EFs, each PFT was linked to a fire biome (see Table A1). INFERNO uses 205

these to estimate emissions (Eq. 9). 206

$$207 E_{X,PFT} = EC_{PFT} EF_{X,PFT} / [C] (9)$$

208 Here E_X is the amount of species X emitted by fires (in kg m⁻² s⁻¹), EC is the emitted carbon (in kgC m⁻²

209 s⁻¹) and EF_X is the PFT-specific emission factor (see Table 1) (in kg of species emitted per kg of biomass

210 burnt), and [C] is the dry biomass carbon content which we assume as 50% (a common simplification;

211 Lamlom and Savidge, 2003). INFERNO currently provides emissions for basic trace gases: CO2, CO,

212 CH₄, NO_x, SO₂ and aerosols: organic carbon (OC) and black carbon (BC).

213 Table 1. INFERNO's emission factors per PFT created from the emission profiles in Akagi et al. (2011), such 214 215 that each PFT was attributed a fire biome (see Suppl. 2). This method of attributing emission factors to PFTs is similar to that presented in Thonicke et al. (2010), and can be extended to include all species of trace gases 216 and aerosols compiled in Akagi et al. (2011).

Emission Factors (g / kg)	CO_2	CO	CH ₄	NO_x	SO_2	OC	BC
Broadleaf Evergreen Tree (Tropical)	1643	93	5.07	2.55	0.40	4.71	0.52
Broadleaf Evergreen Tree (Temperate)	1637	89	3.92	2.51	0.40*	8.2**	0.56**
Broadleaf Deciduous Tree	1643	93	5.07	2.55	0.40	4.71	0.52
Needleleaf Evergreen Tree	1637	89	3.92	2.51	0.40*	8.2**	0.56**
Needleleaf Deciduous Tree	1489	127	5.96	0.90	0.40*	8.2**	0.56**
C3 grass	1637	89	3.92	2.51	0.40*	8.2**	0.56**
C4 grass	1686	63	1.94	3.9	0.48	2.62	0.37
Evergreen Shrub	1637	89	3.92	2.51	0.40*	8.2**	0.56**
Deciduous Shrub	1489	127	5.96	0.90	0.40*	8.2**	0.56**

*Profile not available in Akagi et al. (2011), therefore we mimic tropical forests; **from Andreae and Merlet (2001).

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225 2.2 Implementation within JULES

- INFERNO is currently implemented within the Joint UK Land Environment Simulator (JULES). (Best 226
- 227 et al., 2011; Clark et al., 2011) its carbon fluxes and vegetation dynamics. The results shown here used
- JULES v4.3.1 and INFERNO will be included in JULES from version 4.5 onwards. INFERNO utilizes 228
- 229 soil moisture (see Eq. 6,8), which JULES calculates as the balance between precipitation (following the
- 230 scheme for rainfall interception in (Johannes Dolman and Gregory, 1992)) and extraction by
- 231 evapotranspiration and runoff (Cox et al. 1999; Best et al. 2011). JULES has four soil layers, and 232 INFERNO uses the top layer unfrozen soil moisture (0 to 0.1 m depth). Note that in its current state,
- 233
- JULES does not associate carbon pools with depths, hence it is not possible to access the top-most DPM
- only for example. The vegetation dynamics and litter carbon used obey the TRIFFID DGVM (Cox, 234
- 235 2001). Fractional coverage of PFTs in any gridcell is based on competition for resources (light and
- 236 water), and governed by Lotka-Volterra competition equations, and based on a tree-shrub-grass
- 237 dominance hierarchy (Cox, 2001).
- 238 In JULES, vegetation carbon content is determined by the balance between photosynthesis, respiration,
- 239 and litterfall. Within JULES, TRIFFID (the Top-down Representation of Foliage and Flora Including
- 240 Dynamics; Cox, 2001) predicts changes in biomass and the fractional coverage of nine plant functional
- 241 types (Table A1) based on accumulated carbon fluxes and height-based competition, where the tallest
- 242 trees have the first access to space (Harper et al. In Prep). Vegetation can grow in height, and the carbon
- 243 in leaves, roots, and wood is related allometrically to the "balanced LAI", L_b (Cox_2001). L_b is the
- seasonal maximum leaf area index (LAI) and a function of plant height. Within INFERNO, leaf carbon 244
- ($Leaf_C$, used for calculating FD and emissions) is: 245
- 246 $Leaf_C = \sigma_l L_b$ (10)
- 247 Meanwhile, wood carbon ($Wood_C$, which affects emissions), is calculated as:
- $Wood_C = a_{wl} L_b^{b_{wl}}$ 248 (11)
- 249 PFT dependent parameters (σ_l) , the Specific Leaf Density, a_{wl} , the allometric coefficient and b_{wl} , the
- 250 allometric exponent) are given in Table A1.
- 251 When using JULES in its standalone version, INFERNO can use inputs of population density (in people
- 252 km⁻²) and cloud-to-ground lightning flash rates (in flashes km⁻² month⁻¹) from ancillary datasets.
- 253 Interestingly, lightning can be interactively simulated in atmospheric models (not population), although
- 254 this will not be explored in this paper. Similarly, meteorology needs to be prescribed and is then
- 255 interpolated from its native temporal resolution to the model's time-step. Although designed to be
- integrated within an ESM, the capability to run INFERNO with JULES only is particularly useful for 256
- 257 present-day comparison with observations, and to dissociate causes of biases in results. In its current
- 258 early state, INFERNO provides a diagnostic tool, it does not remove carbon from vegetation nor does it
- 259 lead to tree mortality.

260

2.3 Fire Weather Indices

- 261 Three other well-established daily fire indices are also available within JULES. These indices have been
- 262 used for several decades to help plan operational response to wildfires on Numerical Weather Predictions

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- 266 (NWP) timescales. Although unit-less and ill-defined risk-based quantities, comparison to INFERNO is
- still useful for understanding the results in the context of practically established metrics.
- 268 The Canadian Fire Weather Index (Forestry Canada, 1992; Van Wagner and Pickett, 1985) consists of
- 269 six components, calculated from basic meteorological parameters. Three are fuel moisture codes
- designed to represent the drying of different fuel types, their characteristics are displayed in Table A2.
- 271 Two intermediate quantities, the Initial Spread Index and the build-up index are calculated from these,
- and are in turn used to yield the final Fire Weather Index (FWI):

273
$$FWI = \begin{cases} e^{2.72(0.434 \ln B)^{0.647}}, B > 1 \\ B, B \le 1 \end{cases}$$
 (12)

- where B = 0.1 ISI FD with ISI the Initial Spread Index and FD the Fuel Density. We refer to the
- 275 original publications for detailed equations for the complex Canadian FWI and each of its components.
- 276 The McArthur Forest Fire Danger Index (Noble et al., 1980; Sirakoff, 1985) was developed for use in
- 277 Australia. Simpler in its formulation than the Canadian index, it consists of a drought component
- 278 modified by the local temperature, humidity and wind speed. The calculation of the drought component
- 279 depends on the soil moisture deficit (the amount of water needed to restore the soil moisture content of
- the top 800 mm of soil to 200 mm), which is related to the JULES soil moisture.
- The FFDI $(F_{McArthur})$ is given by:
- $F_{McArthur} = 1.275D^{0.987}e^{\frac{T}{29.5858} \frac{H}{28.9855} + \frac{W}{42.735}}$ (13)
- where T is the daily maximum temperature, H the daily minimum relative humidity and W the daily
- mean wind speed. And D is the drought factor, given by:

285
$$D = \frac{0.191(l+104)(N+1)^{1.5}}{3.25(N+1)^{1.5}+R-1}$$
 (14)

- where N is the number of days since the last rain, R the total rain in the most recent day with rain and I
- the amount of rain needed to restore the soil moisture content to 200 mm in the top 800 mm of soil.
- 288 Finally, the Nesterov Index (Nesterov, 1949) is the simplest fire index implemented in JULES. It uses
- 289 only the daily mean temperature, mean daily dew point (or suitable substitute), daily total precipitation
- and the previous day's index. The index is incremented daily, unless daily precipitation exceeds 3 mm,
- 291 in which case it is reset:

292
$$N = \begin{cases} N_0 + T(T-D), P < 3mm \\ 0, P \ge 3mm \end{cases}$$
 (15)

- where T is the mean daily temperature, D the mean daily dewpoint, P the daily total precipitation and
- 294 N₀ the previous day's index. The Nesterov index is a key component for other fire models (Venevsky et
- 295 al., 2002; Thonicke et al., 2010)

3 Model configuration

296

- 297 Monthly lightning data was obtained from LIS-OTD (Lightning Imaging Sensor-Optical Transient
- 298 Detector) observations for 2013 (Christian et al., 2003) and was recycled for every year in the simulation.
- 299 These detections were converted to cloud-to-ground strikes using the relationship presented in (Prentice
- 300 and Mackerras, 1977). Land use and population density were obtained from the HYDE dataset (Hurtt et
- 301 al., 2011) and then linearly interpolated to create inter-annually varying data. Finally annual CO₂

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306 concentrations, which affect vegetation dynamics, were prescribed as a global average following the 307 dataset prepared for the global carbon budget (Le Quéré et al., 2015). To test the sensitivity to the meteorological input, JULES simulations were driven by meteorology from 308 both CRU-NCEP (Climate Research Unit and -National Center for Environmental Prediction) v5 309 (http://dods.extra.cea.fr/data/p529viov/cruncep/), and WFDEI (Weedon et al., 2014) with precipitation 310 from the GPCC (Schneider et al., 2013). Both datasets were used on a 6-hourly basis. 311 312 Outside of these driving variables, JULES was configured according to the TRENDY project (Sitch et 313 al., 2015)(Peng et al., 2015)(Peng et al., 2015). 100 year spin-up was performed repeating the 1990-2000 conditions tenfold. Four configurations were used to create simulations covering 1990-2013, although to 314 315 validate INFERNO only the 1997-2010 period was analysed. The first three use CRU-NCEP 316 meteorology with each of our three ignitions modes (see Sect. 2.1.1); constant ignitions (mode 1), 317 prescribed lightning and constant anthropogenic ignitions (mode 2), and both natural and anthropogenic 318 ignitions varying with prescribed lightning and population density (mode 3). The fourth simulation 319 assumes mode 1 (constant ignitions), while meteorology is prescribed from WFDEI and precipitation 320 321 Evaluation was performed against the published data for GFEDv3, FINNv1, GFAS and GFEDv4. We 322 also used the data from GFEDv4s (http://globalfiredata.org, manuscript in preparation) and GFEDv4 323 (Giglio et al., 2013) to calculate grid-specific emissions and burnt-area. The Global Fire Emissions 324 Database (GFED) passes satellite observation of burnt area through the Carnegie-Ames-Stanford-325 Approach (CASA) biogeochemical model in order to obtain emissions from open burning. GFEDv4 326 (Giglio et al., 2013) innovates on GFEDv3 (Giglio et al., 2010) mainly through an updated algorithm to 327 retrieve burnt area from MODIS satellite products and an increased spatial and temporal resolution, to 328 0.25° and daily (this resolution was assessed in Mangeon et al., 2015). Meanwhile GFEDv4s also 329 includes the contribution from small fires (Randerson et al., 2012). The Fire Inventory from NCAR 330 version 1.0 (FINNv1, Wiedinmyer et al., 2011) provides high-resolution (both temporal and spatial) 331 global emissions of trace gas and particle emissions from open burning of biomass. It focuses on rapid 332 availability and assimilation in real time forecast and follows a similar process to GFED to estimate 333 emission, but its burnt area is obtained directly from fire pixel using land cover (Wiedinmyer et al.,

4 Results

334

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337

Maps of the burnt area and emitted carbon are displayed in Fig. 2, their resolution is 192 longitudes by
145 latitudes grid-cells (1.875°x1.24°). The results from INFERNO used a configuration with CRUNCEP
meteorology and the third ignition mode: interactive lighting and anthropogenic ignitions. We compare
our results with downscaled means from GFED. INFERNO accurately diagnoses total fire occurrence
and emissions over the 1997-2010 period: we found a spatial correlation of R=0.66 for burnt area and
R=0.59 for emitted carbon, both passing the t-test with 95% significance. In addition, regional mean
yearly budgets are compared with GFED in Table B1. Compared to GFEDv4, we notice INFERNO

2011). The Global Fire Assimilation System (GFAS, Kaiser et al., 2012), unlike the aforementioned

products, directly assess emissions from satellite-observed fire radiative power more apt at detecting

small fires and avoiding the uncertainty of biogeochemical models.

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estimates higher burnt area in all regions apart from Australia and New Zealand, and southern hemisphere Africa. Meanwhile emitted carbon is underestimated in boreal regions and equatorial Asia, but overestimated in most other regions (significantly in southern hemisphere America). Over the studied period, C4 grass were the main contributors to burnt area in INFERNO (a mean 2.34 Mkm² per year), meanwhile Broadleaf Evergreen Trees (Tropical) led to the most emitted carbon (a mean 1.48 Pg per year). GFEDv4 projects the grid-box with maximum burnt area within the Central African Republic (87% of grid fraction burnt per year), while INFERNO finds a maximum burnt area of 57%, slightly to the North (south-east of lake Tchad). The discrepancy is much larger for emissions, with a maximum emitted carbon of 1.47 kg per m² in Indonesia predicted by GFEDv4s, against 0.4 kg per m² for INFERNO, in Angola. These results could be expected, as INFERNO focuses on capturing global biomass burning, it will not represent such extremes of burning, furthermore the immense emitted carbon observed in Indonesia follows from undiagnosed peat fires, INFERNO's approach to burnt area only considers trees, grass and shrub cover and was determined heuristically, meanwhile Hantson et al. (2014) found global fire size was mostly influenced by precipitation, aridity and human activity (population density and croplands). Further parameterizations for fire size exist (e.g., Hantson et al., 2015, 2016) which could improve INFERNO burnt area estimates while maintaining simplicity and traceability.

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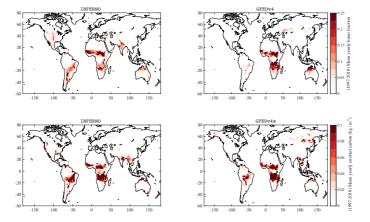


Fig. 2. 1997-2010 mean yearly burnt fraction (above) and emitted carbon (below, in kg m⁻²). Shown for INFERNO on the left (with CRUNCEP meteorology and interactive ignitions (mode 3) and for GFED on the right.

Figure 3 shows the modelled global annual average biomass burning emissions and burnt area from 1997 to 2010. The three ignition methods are evaluated: fully interactive ignitions (red) predict the highest carbon emissions while interactive lightning with constant human ignitions (blue) the lowest. WFDEI was observed to lead to more biomass burning emissions in tropical forest areas (and in particular the borders of rainforests), while CRU-NCEP favoured burning in near-desert areas (the Sahel, India and south American grasslands). We expect this result to be significantly influenced by differences in precipitation (GPCC for WFDEI runs and CRU for CRU-NCEP; Schneider et al., 2013).

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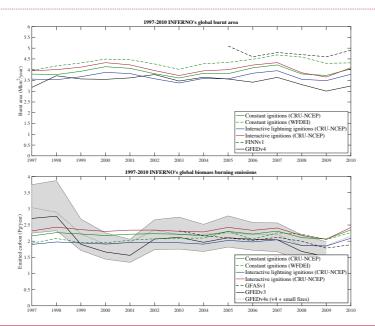
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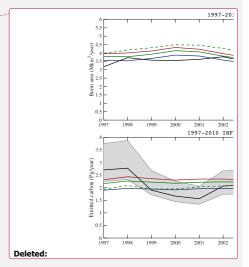


Fig. 3. 1997-2010 biomass burning emissions and burnt area predicted by INFERNO. Two driving datasets were used, CRU-NCEP (solid lines) and WFDEI (green dotted line). Observations are shown in black (MODIS-based estimates). The grey shading represents one standard deviation within GFEDv3's estimates.

Comparisons to FINNv1, GFEDv4, GFASv1 and GFEDv3 were restricted to their budgets published in Kaiser et al. (2012), van der Werf et al. (2010), Wiedinmyer et al. (2011) and Giglio et al. (2013) respectively. Meanwhile we calculated global emissions from GFEDv4s (http://globalfiredata.org, manuscript in preparation).

Biomass burning emissions and burnt area simulated by the model follow similar trends to GFEDv3, although with a smaller inter-annual variability in the model. Carbon emissions from all simulations fall within one standard deviation of GFEDv3, apart from three years: 1997, 1998 and 2001. Note that for these years, emissions in GFED were obtained from the lower resolution AVHRR rather than MODIS. 1997 and 1998 were strong El-Niño years during which droughts in equatorial Asia led to extreme emissions from land-clearing fires, a recurrent problem in the region (Field et al., 2009). Indeed in 1997, in the region contained between 20S-20N and 90E-160E (or equatorial Asia), GFEDv3 estimate emissions of 1.07 PgC, while INFERNO (with CRU-NCEP and fully interactive ignitions) estimates 0.15 PgC. Unfortunately, peat is not modelled in JULES and thus neither is peat present in our fire scheme. It was estimated tropical peat fires alone produced an average of 0.1 PgC per year for 1997-2009, and 0.7 PgC in 1997 in particular (van der Werf et al., 2010). Furthermore, 2002 and 2006 also saw important peat burning, with GFEDv3 estimating peat emissions of 0.16 and 0.21 PgC respectively. In both of these years, the trend in INFERNO differs from GFEDv3's (stagnation in 2002 and decrease in 2006). Peat-lands can be significant in equatorial Asia but also boreal regions where their combustion

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leads to the release of long-stored carbon (Turetsky et al., 2015). In 1998 and 2001, the difference in emissions could not be attributed to a particular location. While fire emissions from Equatorial Asia were underestimated, GFEDv3 observed lower emissions over Africa compared to INFERNO, which seems to be the key driver of our discrepancies.

Table 2. Mean yearly emission budgets in Peta-grams of emitted carbon and mean yearly burnt area budgets in Mkm2 for the 1997-2010 period. Latitudes were bound to: beyond 50° (high latitudes), 35° to 50° (mid-latitudes), 15° to 35° (low latitudes) and below 15° (equatorial). Four configurations of INFERNO are presented, with CRU-NCEP and WFDEI driving meteorology coupled with three ignition modes: mode 1 indicates constant anthropogenic and lightning ignitions, mode 2 is for constant anthropogenic with interactive lightning ignitions, and mode 3 for interactive lightning and anthropogenic ignitions.

Emitted carbon (PgC/year)	mode 1 CRU-NCEP	mode 1 WFDEI	mode 2 CRU-NCEP	mode 3 CRU-NCEP
High latitudes	0.087	0.096	0.082	0.091
Mid-latitudes	0.185	0.193	0.170	0.191
Low latitudes	0.716	0.624	0.627	0.591
Equatorial	1.157	1.130	1.021	1.385

Burnt area (Mkm² / year)	mode 1 CRU-NCEP	mode 1 WFDEI	mode 2 CRU-NCEP	mode 3 CRU-NCEP
High latitudes	0.176	0.196	0.162	0.179
Mid-latitudes	0.485	0.557	0.445	0.531
Low latitudes	1.648	1.884	1.558	1.531
Equatorial	1.524	1.580	1.423	1.693

Table 2 shows the budgets for four latitudinal bands across the various simulations performed. The second ignition mode (constant anthropogenic and interactive lightning ignitions at any time and place) appears to consistently predict lower emissions and burnt area (with the exception of low latitudes). Furthermore, the main impact of using an ignition model that varies with both natural and anthropogenic ignitions is a reduction of fires at low (tropical and sub-tropical) latitudes, and an increase in equatorial regions. Indeed, when compared to constant ignitions (mode 1), interactive ignitions (mode 3) predict more emissions in forest encroachment regions (noticeably surrounding the Congo and Amazon rainforests), and less in heavily-populated areas (Nigeria, India). Meanwhile, we observed interactive lightning ignitions (mode 2) significantly reduced burning in grassland-savannah environments. We link this to the predominance of cloud-to-ground lightning strikes in wet environment within the LIS-OTD dataset (e.g. the Congo rainforest, (Christian et al., 2003) and fewer strikes (and ignitions) in the more flammable grasslands and savannahs. These issues are visible in Fig. B1, which shows difference maps of the four model configurations, for 1997-2010 mean yearly totals. Equatorial and boreal regions include

peat that leads to large fuel consumption, which is unaccounted for in JULES, suggesting that our model will inherently underestimate emissions from these regions.

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INFERNO

GFEDv3

FINNv1

Species-specific average emissions produced by the INFERNO scheme are shown in Table 3 in Tg per year for the 1997-2010 period. CO and CH4 appear to be produced in noticeably larger quantities than in observation-based emission estimates. This hints at an overrepresentation of smouldering-type combustion. In INFERNO this might be due to the emission factors used, or the type of vegetation burnt.

Table 3. Average annual emission (Tg per year) for INFERNO with the interactive ignition mode and CRUNCEP reanalysis (3 – CRUNCEP) and the constant ignition mode and WFDEI reanalysis (1 – WFDEI), comparison to GFASv1 (Kaiser et al., 2012), GFEDv3 (van der Werf et al., 2010) and FINNv1 (Wiedinmyer et al., 2011) is provided.

Global emission (Tg/year)	<u>CO₂</u>	CO	<u>CH</u> ₄	NO _x	<u>BC</u>	<u>oc</u>	- -	Formatteu rabie
3 - CRUNCEP.	7510.7	455.5	26.5	12.8	2.6	26.3	- "	Formatted: Subscript Formatted: Subscript
RNO 1 – WFDEI	7149.8	429.3	24.8	12.2	2.4	24.9		Formatted: Subscript Formatted: Font:Bold
GFASv1	6906.7	351.5	19.0	9.5	2.0	18.2	The same of the sa	Formatted: Font:Bold
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In order to examine whether our flammability can represent fire occurrence, three other fire indices were diagnosed, namely the McArthur, Nesterov and Canadian fire indices. These indices were obtained seamlessly during the model runs, therefore utilizing the same meteorological and hydrological driving variables, and the same vegetation conditions. Their predictions were regressed with GFEDv4 1997-2010 annual burnt area (Giglio et al., 2013). This analysis relies on the assumption that fire indices can be used as a proxy for the variability of fire occurrence and spread, and eventually of burnt area (not the magnitude). Only areas that had been observed to burn sometime between 1997 and 2010 were sampled; to avoid accounting for high fire indices in non-vegetated areas such as the Sahara.

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372.5

15.7

18.2

9.4

12.5

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7322.8

Table 4 shows the result of our analysis. Ignitions followed mode 1; in this mode ignitions are constant, therefore the only variability in burnt area (and performance) is due to INFERNO's flammability scheme.

The McArthur index performs poorly at high latitudes (it was made for Australia), but outperforms the other indices in low latitude regions. The Canadian and Nesterov indices correlate best with observed burnt area in high latitude regions (for which they were developed). Altogether, INFERNO's burnt area appears to follow observed burnt area better than the sole usage of a fire index.

Table 4. Temporal correlation coefficients (R) of annual means (1997-2010) shown for four latitudinal bands. R-coefficients were obtained between either of the three simulated fire indices or INFERNO's burnt area (ubiquitous ignitions – ignition mode I, using CRU-NCEP meteorology) and burnt area from GFEDv4 (Giglio et al., 2013). Italics mean the correlation was not significant (p-value above 0.05). We restrict our analysis to grid-boxes in which GFEDv4 observed burning. Latitudes were bound to: beyond 50° (high latitudes), 35° to 50° (mid-latitudes), 15° to 35° (low latitudes) and below 15° (equatorial).

R-coefficient (with GFEDv4 burnt area)	INFERNO Burnt area		McArthur Index	Canadian Index	 	Deleted:
Global	0.649	0.088	-0.009	0.266		Formatted: Font:Italic

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High latitudes	0.476	0.522	.0.005	0.519
Mid-latitudes	0.179	_0.006	0.069	0.060
Low latitudes	0.603	0.476	0.499	0.480
Equatorial	0.689	0.239	0.354	0.392

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5 Conclusion

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Through a minimalistic approach we propose a parameterization for fire occurrence of appropriate complexity for application at large spatial scales within an ESM context: the INteractive Fire and Emission algoRithm for Natural envirOnments (INFERNO). It directly only varies according to precipitation (and resulting soil moisture), temperature and humidity, and indirectly it utilizes vegetation. It is also capable of explicitly simulating ignitions using lightning and anthropogenic information. While our scheme manages to represent fire occurrence on large scales (both spatial and temporal), it performs best at low latitudes. INFERNO's burnt area scheme appears superior to the use of fire indices alone (Nesteroy, McArthur and basic Canadian) for capturing annual burnt area variations, and thus one form of fire impact. However, due to the nature of our analysis (fire danger and burnt area remain different quantities) this does not imply INFERNO should supersede fire weather indices for operational purposes, neither has our algorithm been built for numerical weather prediction or seasonal fire danger forecasting. Nonetheless, our current simulations suggest the variability in emissions is underestimated by INFERNO, in particular the impact of the 1997-1998 El-Niño and the subsequent La Niña, which may be attributable to the lack of representation of peat in the model, critical to biomass burning in equatorial Asia and boreal areas. The use of different present-day meteorological datasets has an important impact on the magnitude and variability of our diagnostics. Using WFDEI-GPCC rather than CRU-NCEP led to more burnt area but lower fuel consumption and eventually less emitted carbon (this follows from grasslands burning rather than forests). Vegetation zone interfaces were key to this difference. Similarly, lightning appears to more frequently ignite fires in wet environments (rainforests) while flammable environments (savannah, grasslands) with rarer lightning are sensitive to the presence of an anthropogenic ignition source. Including a scheme to parameterise human impacts appears to significantly reduce fires in heavily populated areas, while favouring their encroachment of rainforests (the vicinity of which are an anthropogenic ignition 'sweet spot' in our parameterization). Nevertheless there is much uncertainty attributed to human induced emissions and effects on fire regime (Marlon et al., 2008; Thonicke et al., 2010). Accordingly, we include different modes to examine the impact of ignitions (human or natural) in INFERNO The implementation of INFERNO within the Met Office's Unified Model and its significance for

present-day atmospheric composition and climate will be investigated in a separate paper. To close the

vegetation-fire feedback, INFERNO will eventually need to remove carbon from vegetation and to

include tree mortality. While a strength of the model is its minimalistic approach the scheme holds

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potential for improvements For instance, litter influences flammability but only live vegetation leads to emissions while in reality litter significantly contributes to observed fuel consumption (van Leeuwen et al., 2014). Similarly, we predict that the inclusion of peat within JULES would improve its fire diagnostics, especially for locations with large fuel consumptions (e.g. equatorial Asia and boreal climates; van der Werf et al., 2010). Given the predictability of emissions from peat fires in relation with precipitation (van der Werf et al., 2008), this would be a promising area of exploration. The value of this model being its simplicity and linearity, any improvements to INFERNO should follow this vision; complex parameterizations are better suited for process-based fire schemes (e.g., Lasslop et al., 2014; Li et al., 2013, p.1).

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Code availability

Information on the JULES land surface model can be found at: http://jules-lsm.github.io/. INFERNO is included in JULES vn4.5 and is included in this documentation. The JULES source code can be accessed via the Met Office's science repository (requires registration): https://code.metoffice.gov.uk/trac/jules. In particular, the version of the code used to produce the outputs included in this study can be accessed at:

https://code.metoffice.gov.uk/trac/jules/browser/main/branches/dev/stephanemangeon/vn4.3.1_inferno.

Appendix A

This appendix contains additional information relating to the INFERNO scheme.

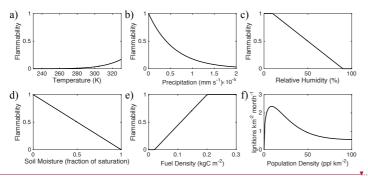
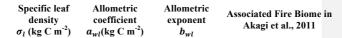
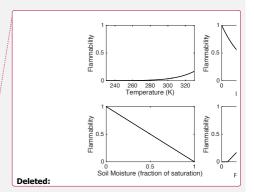


Fig. A1. The <u>mathematical functions used for individual dependencies of INFERNO on key driving variables for flammability (a,b,c,d,e) and ignitions (f), within the range of reasonable earth observations.</u> Note the population density only influences the model output if ignition mode 3 is selected (interactive lightning and human ignition).

Table A1. The key JULES PFT-specific parameters for allometry and vegetation carbon used in our simulations (Clark et al., 2011).





Broadleaf Evergreen Tree (Tropical)	0.0375	0.65	1.667	Tropical Forests
Broadleaf Evergreen Tree (Temperate)	0.0375	0.65	1.667	Temperate Forests
Broadleaf Deciduous Tree	0.0375	0.65	1.667	Tropical Forests
Needleleaf Evergreen Tree	0.1	0.65	1.667	Temperate Forests
Needleleaf Deciduous Tree	0.1	0.75	1.667	Boreal Forests
C3 grass	0.025	0.005	1.667	Temperate Forests
C4 grass	0.05	0.005	1.667	Savannah and Grasslands
Evergreen Shrub	0.05	0.10	1.667	Temperate Forests
Deciduous Shrub	0.05	0.10	1.667	Boreal Forests

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Table A2. The characteristics of the Canadian's Fire Weather Index's three fuel moisture codes.

	Type of fuel	Dry weight (kg m ⁻²)	Time lag (days)	Water capacity (mm)
Fine Fuel Moisture Code	Litter and other fine fuels	0.25	2-3	0.6
Duff Moisture Code	Loosely compacted decomposing organic matter	5	12	15
Drought Code	Deep layer of compact organic matter	25	52	100

559 Appendix B

 $560 \qquad \text{This appendix contains additional results illustrating the dependence of INFERNO with ignitions and its} \\$

performance on a regional basis.

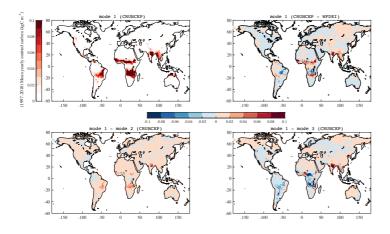


Fig. B1. Emitted carbon difference maps between the four runs performed to analyse the sensitivity of INFERNO to ignitions (our three ignition modes, see Sect. 2.1.1) and meteorology (CRUNCEP and WFDEI-GPCC).

Table B1. Regional budgets according to the standard GFED regions (van der Werf et al., 2010).

		ly Burnt Area Mha)	Mean Yearly Emitted Carbon (TgC)		
GFED standard regions	GFED <u>v</u> 4*	INFERNO**	GFED3***	INFERNO**	
Boreal North America	2.2	5.2	54	37	
Temperate North America	1.8	29.9	9	106	
Central America	1.8	7.9	20	45	
Northern Hemisphere South America	2.6	4.0	22	51	
Southern Hemisphere South America	18.7	68.3	271	483	
Europe	0.7	5.0	4	29	
Middle East	0.8	12.3	2	19	
Northern Hemisphere Africa	117.7	120.4	481	533	
Southern Hemisphere Africa	125.0	57.6	557	610	
Boreal Asia	5.6	9.7	128	55	
Central Asia	13.6	23.8	36	50	
Southeast Asia	7.0	29.6	103	170	

Equatorial Asia	1.6	0.5	191	10
Australia and New Zealand	50.2	30.2	135	96

* GFED_v4 mean yearly burnt area from Giglio et al. (2013), from 1997 to 2011. ** INFERNO mean yearly burnt area from 1997 to 2010, using ignition mode 3 (varying anthropogenic and natural ignitions) and CRU-NCEP driving meteorology. *** GFED3 mean yearly emitted carbon from van der Werf et al. (2010) from 1997 to 2009.

Author contribution

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- Apostolos Voulgarakis supervised the scientific design of INFERNO and the writing of this article. Gerd 571
- 572 Folberth also supervised these aspects, with an emphasis on technical aspects of INFERNO in relation
- 573 with the Met Office's Unified Model. Richard Gilham contributed to the technical design of the model
- 574 and its implementation and led the writing on fire indices. Anna Harper contributed to the design of
- 575 INFERNO in relation to the vegetation scheme's recent development, helped with the analysis of
- vegetation biases in the study's results and led the writing on the vegetation scheme. Stephen Sitch 576
- contributed throughout the writing, analysis and the scientific design of this study. 577

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585 References

- 586 Akagi, S. K., Yokelson, R. J., Wiedinmyer, C., Alvarado, M. J., Reid, J. S., Karl, T., Crounse, J. D. and
- 587 Wennberg, P. O.: Emission factors for open and domestic biomass burning for use in atmospheric
- 588 models, Atmos Chem Phys, 11(9), 4039-4072, doi:10.5194/acp-11-4039-2011, 2011.
- Arora, V. K. and Boer, G. J.: Fire as an interactive component of dynamic vegetation models, J. Geophys. 589
- 590 Res. Biogeosciences, 110(G2), G02008, doi:10.1029/2005JG000042, 2005.
- 591 Best, M. J., Pryor, M., Clark, D. B., Rooney, G. G., Essery, R. . L. H., Ménard, C. B., Edwards, J. M.,
- 592 Hendry, M. A., Porson, A., Gedney, N., Mercado, L. M., Sitch, S., Blyth, E., Boucher, O., Cox, P. M.,
- 593 Grimmond, C. S. B. and Harding, R. J.: The Joint UK Land Environment Simulator (JULES), model
- 594 description - Part 1: Energy and water fluxes, Geosci. Model Dev., 4(3), 677-699, doi:10.5194/gmd-4-
- 595 677-2011, 2011.
- Bond, W. J. and Keeley, J. E.: Fire as a global "herbivore": the ecology and evolution of flammable 596
- 597 ecosystems, Trends Ecol. Evol., 20(7), 387-394, doi:10.1016/j.tree.2005.04.025, 2005.
- Bowman, D. M., Murphy, B. P., Boer, M. M., Bradstock, R. A., Cary, G. J., Cochrane, M. A., Fensham, 598
- 599 R. J., Krawchuk, M. A., Price, O. F. and Williams, R. J.: Forest fire management, climate change, and
- 600 the risk of catastrophic carbon losses, Front. Ecol. Environ., 11(2), 66-67, doi:10.1890/13.WB.005,
- 601
- 602 Bowman, D. M. J. S., Balch, J. K., Artaxo, P., Bond, W. J., Carlson, J. M., Cochrane, M. A., D'Antonio,
- 603 C. M., DeFries, R. S., Doyle, J. C., Harrison, S. P., Johnston, F. H., Keeley, J. E., Krawchuk, M. A., Kull,
- 604
- C. A., Marston, J. B., Moritz, M. A., Prentice, I. C., Roos, C. I., Scott, A. C., Swetnam, T. W., Werf, G. R. van der and Pyne, S. J.: Fire in the Earth System, Science, 324(5926), 481–484, 605
- doi:10.1126/science.1163886, 2009. 606
- Christian, H. J., Blakeslee, R. J., Boccippio, D. J., Boeck, W. L., Buechler, D. E., Driscoll, K. T., 607
- 608 Goodman, S. J., Hall, J. M., Koshak, W. J., Mach, D. M. and Stewart, M. F.: Global frequency and

- 609 distribution of lightning as observed from space by the Optical Transient Detector, J. Geophys. Res.
- 610 Atmospheres, 108(D1), 4005, doi:10.1029/2002JD002347, 2003.
- Chuvieco, E., Giglio, L. and Justice, C.: Global characterization of fire activity: toward defining fire
- regimes from Earth observation data, Glob. Change Biol., 14(7), 1488-1502, 2008. 612
- 613 Clark, D. B., Mercado, L. M., Sitch, S., Jones, C. D., Gedney, N., Best, M. J., Pryor, M., Rooney, G. G.,
- Essery, R. L. H., Blyth, E., Boucher, O., Harding, R. J., Huntingford, C. and Cox, P. M.: The Joint UK Land Environment Simulator (JULES), model description Part 2: Carbon fluxes and vegetation dynamics, Geosci Model Dev, 4(3), 701–722, doi:10.5194/gmd-4-701-2011, 2011. 614
- 615
- 616
- Cox, P. M.: Description of the TRIFFID dynamic global vegetation model, Technical Note 24, Hadley
- Centre, United Kingdom Meteorological Office, Bracknell, UK. [online] Available from: http://www.metoffice.gov.uk/media/pdf/9/h/HCTN_24.pdf (Accessed 10 September 2015), 2001. 618
- 619
- 620 Field, R. D., van der Werf, G. R. and Shen, S. S. P.: Human amplification of drought-induced biomass
- burning in Indonesia since 1960, Nat. Geosci., 2(3), 185-188, doi:10.1038/ngeo443, 2009. 621
- 622 Forestry Canada: Development and structure of the Canadian Forest Fire Behavior Prediction System.
- [online] Available from: http://cfs.nrcan.gc.ca/publications?id=10068 (Accessed 8 January 2016), 1992. 623
- 624 Fosberg, M. A., Cramer, W., Brovkin, V., Fleming, R., Gardner, R., Gill, A. M., Goldammer, J. G.,
- 625 Keane, R., Koehler, P., Lenihan, J., Neilson, R., Sitch, S., Thornicke, K., Venevski, S., Weber, M. G.
- 626 and Wittenberg, U.: Strategy for a Fire Module in Dynamic Global Vegetation Models, Int. J. Wildland
- 627 Fire, 9(1), 79-84, 1999.
- 628 Giglio, L., Randerson, J. T., van der Werf, G. R., Kasibhatla, P. S., Collatz, G. J., Morton, D. C., and
- 629 DeFries, R. S.: Assessing variability and long-term trends in burned area by merging multiple satellite
- 630 fire products, Biogeosciences, 7, 1171-1186, doi:10.5194/bg-7-1171-2010, 2010.
- 631 Giglio, L., Randerson, J. T. and van der Werf, G. R.: Analysis of daily, monthly, and annual burned area
- 632 using the fourth-generation global fire emissions database (GFED4), J. Geophys. Res. Biogeosciences,
- 118(1), 317–328, doi:10.1002/jgrg.20042, 2013. 633
- Hantson, S., Pueyo, S. and Chuvieco, E.: Global fire size distribution is driven by human impact and 634
- 635 climate, Glob. Ecol. Biogeogr., n/a-n/a, doi:10.1111/geb.12246, 2014.
- 636 Hantson, S., Pueyo, S. and Chuvieco, E.: Global fire size distribution: from power law to log-637 normal, International Journal of Wildland Fire 25, 403-412, doi:10.1071/WF15108, 2016.
- 638 Hantson, S., Lasslop, G., Kloster, S. and Chuvieco, E.: Anthropogenic effects on global mean fire
- 639 size, International Journal of Wildland Fire 24, 589-596, doi:10.1071/WF14208, 2015.
- 640 Harper, A., Cox, P., Friedlingstein, P., Wiltshire, A., Jones, C., Sitch, S., Mercado, L. M., Groenendijk,
- 641 M., Robertson, E., Kattge, J., Bönisch, G., Atkin, O. K., Bahn, M., Cornelissen, J., Niinemets, Ü.,
- Onipchenko, V., Peñuelas, J., Poorter, L., Reich, P. B., Soudzilovskaia, N., and van Bodegom, P.: 642
- 643 Improved representation of plant functional types and physiology in the Joint UK Land Environment
- 644 Simulator (JULES v4.2) using plant trait information, Geosci. Model Dev. Discuss., doi:10.5194/gmd-
- 645 2016-22, in review, 2016.
- 646 Huntrieser, H., Schumann, U., Schlager, H., Höller, H., Giez, A., Betz, H.-D., Brunner, D., Forster, C.,
- O. Pinto Jr. and Calheiros, R.: Lightning activity in Brazilian thunderstorms during TROCCINOX: 647
- implications for NOx production, Atmos Chem Phys Discuss, 7(5), 14813-14894, doi:10.5194/acpd-7-648
- 14813-2007, 2007. 649
- Hurtt, G. C., Chini, L. P., Frolking, S., Betts, R. A., Feddema, J., Fischer, G., Fisk, J. P., Hibbard, K.,
- 651 Houghton, R. A., Janetos, A., Jones, C. D., Kindermann, G., Kinoshita, T., Goldewijk, K. K., Riahi, K.,
- Shevliakova, E., Smith, S., Stehfest, E., Thomson, A., Thornton, P., Vuuren, D. P. van and Wang, Y. P.: 652
- 653 Harmonization of land-use scenarios for the period 1500-2100: 600 years of global gridded annual land-

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- 654 use transitions, wood harvest, and resulting secondary lands, Clim. Change, 109(1-2), 117-161,
- 655 doi:10.1007/s10584-011-0153-2, 2011.
- 656 Johannes Dolman, A. and Gregory, D.: The Parametrization of Rainfall Interception In GCMs, Q. J. R.
 - Meteorol. Soc., 118(505), 455–467, doi:10.1002/qj.49711850504, 1992.
- 658 Johnston, F. H., Henderson, S. B., Chen, Y., Randerson, J. T., Marlier, M., DeFries, R. S., Kinney, P.,
- Bowman, D. M. J. S. and Brauer, M.: Estimated Global Mortality Attributable to Smoke from Landscape 659
- Fires, Environ. Health Perspect., 120(5), 695-701, doi:10.1289/ehp.1104422, 2012. 660
- 661 Kaiser, J. W., Heil, A., Andreae, M. O., Benedetti, A., Chubarova, N., Jones, L., Morcrette, J.-J.,
- 662 Razinger, M., Schultz, M. G., Suttie, M. and van der Werf, G. R.: Biomass burning emissions estimated
- 663 with a global fire assimilation system based on observed fire radiative power, Biogeosciences, 9(1), 527-
- 664 554, doi:10.5194/bg-9-527-2012, 2012.
- 665 Lamarque, J.-F., Bond, T. C., Eyring, V., Granier, C., Heil, A., Klimont, Z., Lee, D., Liousse, C.,
- Mieville, A., Owen, B. and others: Historical (1850-2000) gridded anthropogenic and biomass burning 666
- 667 emissions of reactive gases and aerosols: methodology and application, Atmospheric Chem. Phys.,
- 668 10(15), 7017–7039, 2010.
- 669 Lamlom, S. H. and Savidge, R. A.: A reassessment of carbon content in wood: variation within and
- 670 between 41 North American species, Biomass Bioenergy, 25(4), 381-388, doi:10.1016/S0961-
- 9534(03)00033-3, 2003. 671
- Lasslop, G., Thonicke, K. and Kloster, S.: SPITFIRE within the MPI Earth system model: Model 672
- 673 development and evaluation, J. Adv. Model. Earth Syst., n/a-n/a, doi:10.1002/2013MS000284, 2014.
- Lasslop, G., Hantson, S. and Kloster, S.: Influence of wind speed on the global variability of burned
- 675 fraction: a global fire model's perspective, Int. J. Wildland Fire, 24(7), 989-1000, 2015.
- 676 van Leeuwen, T. T., van der Werf, G. R., Hoffmann, A. A., Detmers, R. G., Rücker, G., French, N. H.
- F., Archibald, S., Carvalho Jr., J. A., Cook, G. D., de Groot, W. J., Hély, C., Kasischke, E. S., Kloster, 677 678
- S., McCarty, J. L., Pettinari, M. L., Savadogo, P., Alvarado, E. C., Boschetti, L., Manuri, S., Meyer, C. P., Siegert, F., Trollope, L. A., and Trollope, W. S. W.: Biomass burning fuel consumption rates: a field 679
- 680 measurement database, Biogeosciences, 11, 7305-7329, doi:10.5194/bg-11-7305-2014, 2014.
- 681 Le Quéré, C., Moriarty, R., Andrew, R. M., Peters, G. P., Ciais, P., Friedlingstein, P., Jones, S. D., Sitch,
- 682 S., Tans, P., Arneth, A., Boden, T. A., Bopp, L., Bozec, Y., Canadell, J. G., Chini, L. P., Chevallier, F., 683
- Cosca, C. E., Harris, I., Hoppema, M., Houghton, R. A., House, J. I., Jain, A. K., Johannessen, T., Kato, 684 E., Keeling, R. F., Kitidis, V., Klein Goldewijk, K., Koven, C., Landa, C. S., Landschützer, P., Lenton,
- 685
- A., Lima, I. D., Marland, G., Mathis, J. T., Metzl, N., Nojiri, Y., Olsen, A., Ono, T., Peng, S., Peters, W., 686
- Pfeil, B., Poulter, B., Raupach, M. R., Regnier, P., Rödenbeck, C., Saito, S., Salisbury, J. E., Schuster, U., Schwinger, J., Séférian, R., Segschneider, J., Steinhoff, T., Stocker, B. D., Sutton, A. J., Takahashi, 687
- 688 T., Tilbrook, B., van der Werf, G. R., Viovy, N., Wang, Y.-P., Wanninkhof, R., Wiltshire, A. and Zeng,
- N.: Global carbon budget 2014, Earth Syst. Sci. Data, 7(1), 47–85, doi:10.5194/essd-7-47-2015, 2015.
- 690 Li, F., Zeng, X. D. and Levis, S.: A process-based fire parameterization of intermediate complexity in a
- Dynamic Global Vegetation Model, Biogeosciences, 9(7), 2761–2780, doi:10.5194/bg-9-2761-2012, 691
- 692 2012.
- 693 Li, F., Levis, S. and Ward, D. S.: Quantifying the role of fire in the Earth system - Part 1: Improved
- 694 global fire modeling in the Community Earth System Model (CESM1), Biogeosciences, 10(4), 2293-
- 695 2314, doi:10.5194/bg-10-2293-2013, 2013.
- 696 Li, F., Bond-Lamberty, B. and Levis, S.: Quantifying the role of fire in the Earth system - Part 2: Impact
- on the net carbon balance of global terrestrial ecosystems for the 20th century, Biogeosciences, 11(5), 697
- 698 1345-1360, doi:10.5194/bg-11-1345-2014, 2014.

- 699 Mangeon, S., R.D. Field, M. Fromm, C. McHugh, and A. Voulgarakis, 2015: Satellite versus ground-
- 700 based estimates of burned area: A comparison between MODIS based burned area and fire agency reports
- 701 over North America in 2007. Anthropocene Rev., early on-line, doi:10.1177/2053019615588790
- 702 Marlier, M. E., DeFries, R. S., Voulgarakis, A., Kinney, P. L., Randerson, J. T., Shindell, D. T., Chen,
- Y. and Faluvegi, G.: El Nino and health risks from landscape fire emissions in southeast Asia, Nat. Clim. 703
- 704 Change, 3(2), 131-136, doi:10.1038/nclimate1658, 2013.
- 705 Marlon, J. R., Bartlein, P. J., Carcaillet, C., Gavin, D. G., Harrison, S. P., Higuera, P. E., Joos, F., Power,
- M. J. and Prentice, I. C.: Climate and human influences on global biomass burning over the past 706
- two millennia, Nat. Geosci., 1(10), 697-702, doi:10.1038/ngeo313, 2008. 707
- 708 Nesterov, V.: Forest fires and methods of fire risk determination, Russ. Goslesbumizdat Mosc., 1949.
- 709 Noble, I. R., Gill, A. M. and Bary, G. a. V.: McArthur's fire-danger meters expressed as equations, Aust.
- 710 J. Ecol., 5(2), 201–203, doi:10.1111/j.1442-9993.1980.tb01243.x, 1980.
- 711 Pechony, O. and Shindell, D. T.: Fire parameterization on a global scale, J. Geophys. Res. Atmospheres,
- 114(D16), D16115, doi:10.1029/2009JD011927, 2009. 712
- 713 Pechony, O. and Shindell, D. T.: Driving forces of global wildfires over the past millennium and the
- forthcoming century, Proc. Natl. Acad. Sci., doi:10.1073/pnas.1003669107, 2010. 714
- Peng, S., Ciais, P., Chevallier, F., Peylin, P., Cadule, P., Sitch, S., Piao, S., Ahlström, A., Huntingford,
- 716 C., Levy, P., Li, X., Liu, Y., Lomas, M., Poulter, B., Viovy, N., Wang, T., Wang, X., Zaehle, S., Zeng,
- 717 N., Zhao, F. and Zhao, H.: Benchmarking the seasonal cycle of CO2 fluxes simulated by terrestrial
- 718 ecosystem models, Glob. Biogeochem. Cycles, 29(1), 2014GB004931, doi:10.1002/2014GB004931,
- 719 2015.
- Pfeiffer, M., Spessa, A. and Kaplan, J. O.: A model for global biomass burning in preindustrial time:
- 721 LPJ-LMfire (v1.0), Geosci Model Dev, 6(3), 643-685, doi:10.5194/gmd-6-643-2013, 2013.
- Prentice, S. A. and Mackerras, D.: The Ratio of Cloud to Cloud-Ground Lightning Flashes in Thunderstorms, J. Appl. Meteorol., 16(5), 545–550, doi:10.1175/1520-722
- 723
- 0450(1977)016<0545:TROCTC>2.0.CO;2, 1977. 724
- 725 Randerson, J. T., Chen, Y., van der Werf, G. R., Rogers, B. M. and Morton, D. C.: Global burned area
- 726 and biomass burning emissions from small fires, J. Geophys. Res. Biogeosciences, 117(G4), G04012,
- 727 doi:10.1029/2012JG002128, 2012.
- Schneider, U., Becker, A., Finger, P., Meyer-Christoffer, A., Ziese, M. and Rudolf, B.: GPCC's new land
- 729 surface precipitation climatology based on quality-controlled in situ data and its role in quantifying the
- global water cycle, Theor. Appl. Climatol., 115(1-2), 15-40, doi:10.1007/s00704-013-0860-x, 2013.
- 731 Sirakoff, C.: A correction to the equations describing the McArthur forest fire danger meter. Aust. J.
- Ecol., 10(4), 481–481, doi:10.1111/j.1442-9993.1985.tb00909.x, 1985. 732
- 733 Sitch, S., Smith, B., Prentice, I. C., Arneth, A., Bondeau, A., Cramer, W., Kaplan, J. O., Levis, S., Lucht,
- 734 W., Sykes, M. T., Thonicke, K. and Venevsky, S.: Evaluation of ecosystem dynamics, plant geography
- 735 and terrestrial carbon cycling in the LPJ dynamic global vegetation model, Glob. Change Biol., 9(2),
- 736 161-185, doi:10.1046/j.1365-2486.2003.00569.x, 2003.
- Sitch, S., Friedlingstein, P., Gruber, N., Jones, S. D., Murray-Tortarolo, G., Ahlström, A., Doney, S. C.,
- 738 Graven, H., Heinze, C., Huntingford, C., Levis, S., Levy, P. E., Lomas, M., Poulter, B., Viovy, N.,
- 739 Zaehle, S., Zeng, N., Arneth, A., Bonan, G., Bopp, L., Canadell, J. G., Chevallier, F., Ciais, P., Ellis, R.,
- 740 Gloor, M., Peylin, P., Piao, S. L., Le Quéré, C., Smith, B., Zhu, Z. and Myneni, R.: Recent trends and
- 741 drivers of regional sources and sinks of carbon dioxide, Biogeosciences, 12(3), 653-679, doi:10.5194/bg-
- 12-653-2015, 2015.

- 743 Spracklen, D. V., Logan, J. A., Mickley, L. J., Park, R. J., Yevich, R., Westerling, A. L. and Jaffe, D. A.:
- 744 Wildfires drive interannual variability of organic carbon aerosol in the western U.S. in summer, Geophys.
- 745 Res. Lett., 34(16), L16816, doi:10.1029/2007GL030037, 2007.
- 746 Thonicke, K., Venevsky, S., Sitch, S. and Cramer, W.: The role of fire disturbance for global vegetation
- dynamics: coupling fire into a Dynamic Global Vegetation Model, Glob. Ecol. Biogeogr., 10(6), 661–
- 748 677, doi:10.1046/j.1466-822X.2001.00175.x, 2001.
- 749 Thonicke, K., Spessa, A., Prentice, I. C., Harrison, S. P., Dong, L. and Carmona-Moreno, C.: The
- 750 influence of vegetation, fire spread and fire behaviour on biomass burning and trace gas emissions:
- 751 results from a process-based model, Biogeosciences, 7(6), 1991–2011, doi:10.5194/bg-7-1991-2010,
- 752 2010.
- 753 Tosca, M. G., Randerson, J. T. and Zender, C. S.: Global impact of smoke aerosols from landscape fires
- 754 on climate and the Hadley circulation, Atmos Chem Phys, 13(10), 5227–5241, doi:10.5194/acp-13-5227-
- 755 2013, 2013.
- 756 Turetsky, M. R., Benscoter, B., Page, S., Rein, G., van der Werf, G. R. and Watts, A.: Global
- 757 vulnerability of peatlands to fire and carbon loss, Nat. Geosci., 8(1), 11–14, doi:10.1038/ngeo2325, 2015.
- Van Wagner, C. E. and Pickett, T. L.: Equations and FORTRAN program for the Canadian Forest Fire
- 759 Weather Index System. [online] Available from: http://www.cfs.nrcan.gc.ca/publications/?id=19973
- 760 (Accessed 8 January 2016), 1985.
- 761 Venevsky, S., Thonicke, K., Sitch, S. and Cramer, W.: Simulating fire regimes in human-dominated
- 762 ecosystems: Iberian Peninsula case study, Glob. Change Biol., 8(10), 984-998, doi:10.1046/j.1365-
- 763 2486.2002.00528.x, 2002.
- 764 Voulgarakis, A. and Field, R. D.: Fire Influences on Atmospheric Composition, Air Quality and Climate,
- 765 Curr. Pollut. Rep., 1(2), 70–81, doi:10.1007/s40726-015-0007-z, 2015.
- 766 Voulgarakis, A., Savage, N. H., Wild, O., Braesicke, P., Young, P. J., Carver, G. D. and Pyle, J. A.:
- 767 Interannual variability of tropospheric composition: the influence of changes in emissions, meteorology
- 768 and clouds, Atmos Chem Phys, 10(5), 2491–2506, doi:10.5194/acp-10-2491-2010, 2010.
- 769 Voulgarakis, A., Marlier, M. E., Faluvegi, G., Shindell, D. T., Tsigaridis, K. and Mangeon, S.:
- 770 Interannual variability of tropospheric trace gases and aerosols: The role of biomass burning emissions,
- 771 J. Geophys. Res. Atmospheres, 120(14), 7157–7173, doi:10.1002/2014JD022926, 2015.
- 772 Weedon, G. P., Balsamo, G., Bellouin, N., Gomes, S., Best, M. J. and Viterbo, P.: The WFDEI
- 773 meteorological forcing data set: WATCH Forcing Data methodology applied to ERA-Interim reanalysis
- data, Water Resour. Res., 50(9), 7505–7514, doi:10.1002/2014WR015638, 2014.
- van der Werf, G. R., Morton, D. C., DeFries, R. S., Olivier, J. G. J., Kasibhatla, P. S., Jackson, R. B.,
- Collatz, G. J. and Randerson, J. T.: CO2 emissions from forest loss, Nat. Geosci., 2(11), 737-738,
- 777 doi:10.1038/ngeo671, 2009.
- van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Mu, M., Kasibhatla, P. S., Morton, D.
- 779 C., DeFries, R. S., Jin, Y. and van Leeuwen, T. T.: Global fire emissions and the contribution of
- deforestation, savanna, forest, agricultural, and peat fires (1997-2009), Atmos Chem Phys, 10(23),
- 781 11707–11735, doi:10.5194/acp-10-11707-2010, 2010.
- van der Werf, G., Dempewolf, J., Trigg, S. N., Randerson, J. T., Kasibhatla, P. S., Giglio, L., Murdiyarso,
- D., Peters, W., Morton, D. C., Collatz, G. J., Dolman, A. J. and DeFries, R. S.: Climate regulation of fire
- 784 emissions and deforestation in equatorial Asia, Proc. Natl. Acad. Sci., 105(51), 20350–20355,
- 785 doi:10.1073/pnas.0803375105, 2008.
- Wiedinmyer, C., Akagi, S. K., Yokelson, R. J., Emmons, L. K., Al-Saadi, J. A., Orlando, J. J. and Soja,
- 787 A. J.: The Fire INventory from NCAR (FINN): a high resolution global model to estimate the emissions
- 788 from open burning, Geosci. Model Dev., 4(3), 625–641, doi:10.5194/gmd-4-625-2011, 2011.

- Yue, C., Ciais, P., Cadule, P., Thonicke, K., Archibald, S., Poulter, B., Hao, W. M., Hantson, S., Mouillot, F., Friedlingstein, P., Maignan, F. and Viovy, N.: Modelling the role of fires in the terrestrial carbon balance by incorporating SPITFIRE into the global vegetation model ORCHIDEE Part 1: simulating historical global burned area and fire regimes, Geosci Model Dev, 7(6), 2747–2767, doi:10.5194/gmd-7-2747-2014, 2014.
- 790
- 792

Two intermediate quantities, the Initial Spread Index and the build-up index are calculated from these, and are in turn used to yield the final Fire Weather Index.

The McArthur Forest Fire Danger Index (Noble et al., 1980; Sirakoff, 1985) was developed for use in Australia. Simpler in its formulation than the Canadian index, it consists of a drought component modified by the local temperature, humidity and wind speed. The calculation of the drought component depends on the soil moisture deficit (the amount of water needed to restore the soil moisture content of the top 800 mm of soil to 200 mm), which is related to the JULES soil moisture.

Finally, the Nesterov Index (Nesterov, 1949) is the simplest fire index implemented in JULES. It uses only the daily mean temperature, mean daily dew point (or suitable substitute), daily total precipitation and the previous day's index. The index is incremented daily, unless daily precipitation exceeds 3 mm, in which case it is reset. The Nesterov index is a key component for other fire models (Venevsky et al., 2002; Thonicke et al., 2010).