### **Reply to Reviewer 1**

Regional models exhibit large uncertainties in the simulation of secondary organic aerosol (SOA) which have substantial impacts on climate due to aerosol-cloud interactions. This paper reviewed the current Volatility Basis Set (VBS) treatments and investigated the model performances in SOA simulation with a series of scenarios by changing the model configuration in chemical mechanisms and aerosol activation parameterization. Results suggest that simulations with VBS treatments present better agreement with observations compared to the traditional OA method, however, parameters such as the enthalpy of vaporization, percentage of fragmentation and functionalization, and POA emissions can largely influence the result. The paper is well written. I would recommend it to be published after minor revisions.

### **Reply:**

We thank the Reviewer for the comments to improve the presentation of the manuscript. Where applicable, suggestions have been taken into consideration and added to the manuscript. Please see below our point0-by-point replies.

Apparently, the POA emissions play an important role in the simulation of SOA. Better performance is suggested in scenarios with increased POA emission. Does that imply that POA emission is underestimated in current NEI emissions? I would suggest the authors to provide some discussion about that.

Reply: Yes, POA emissions are underestimated in current NEI emissions as POA is assumed to be nonvolatile. In the text, this sentence describes the underprediction in POA emissions: "With the semivolatile POA and FF cases in this study, additional IVOC and SVOC emissions are added as three times of the traditional POA emissions from NEI, to account for missing IVOC and SVOC species in the traditional POA emission inventory."

Page 47: "SSummary" should be "Summary"

Reply: This has been modified.

Page 50: Table 4. Note of "The simulations without the suffix "POA" indicate the cases with nonvolatile default POA emissions" need to be clarified, it should be "The simulations without the suffix "POA" or "FF"".

Reply: This has been modified.

Page 52: Table 6. Poor correlation is suggested in most of cases, implying that some important SOA source is missing, biogenic SOA?

Reply: Yes, as mentioned in the text: "The SOA data from the CalNex campaign only consider contributions from a small number of precursors including biogenic precursors (i.e., isoprene,  $\alpha$ -pinene, and  $\beta$ -caryophyllene), and the anthropogenic precursors (i.e., toluene, polycyclic aromatic hydrocarbons (PAHs) and methyl butenol (MBO))". This discussion has been made clearer to include that this reason also likely contributes to the poor correlation.

Page 53: for CASTNET, the simulated Max 8h O3 is very close to the simulated Max 1h O3, especially in CB6 (41.9 vs 41.8), but the observation doesn't (51.8 vs 47.4). Does that mean the model underestimate the peak value of O3?

Reply: Yes, this is likely to be true. In addition, NMBs and NMEs for Max 1h O3 are higher compared to Max 8h O3, which means that the model is not predicting well the transient peak O3 concentrations.

Page 54: "CB05-25%FF-EM3" present different values in Table 8 and 9, while observation is the same. Please double check.

Reply: Table 8 cases use the Grell-Freitas cumulus parameterization scheme, while Table 9 use the MSKF scheme. This has been made clear in the table headers.

Page 55: Figure 1, "A/OC ratios" should be "OA/OC rations"

Reply: This has been modified.

Page 64: Figure 10, it is very interesting that low CDNC shows at the edge of simulation domain, any explanation about that?

Reply: This is likely due to the fact that there are no boundary conditions for CDNC.

### **Reply to Reviewer 2**

The study is extensive, and is suitable for publication in GMD. I particularly like the summary of existing SOA approaches, centred around the VBS. This is a useful addition to the literature for sure. I request the following issues are addressed prior to publication.

### **Reply:**

We thank the Reviewer for the comments to improve the presentation of the manuscript. Where applicable, suggestions have been taken into consideration and added to the manuscript. We hope that we were able to answer all the reviewer's questions adequately. Please see below our point0-by-point replies.

Page 4 line 80. I'm not sure the commonly held notion of computationally expensive SOA schemes according to the number of products should persist as a general discussion. Most, if not all, SOA models assume equilibrium absorptive mass partitioning which rests on Newton based methods requiring a small number of iterations. Is there a range of % contributions, for example, that display the relative cost of SOA schemes versus, say, the gas phase chemistry?

Reply: The new SOA schemes, such as the VBS, is more computationally expensive in comparison to the "traditional" SOA schemes in 3-D models, such as the Odum 2-product model. From our experience, the SAPRC07 scheme is the most expensive compared to the other gas-phase chemistry schemes due to the number of chemical equations. Therefore, the coupled SAPRC07 with the VBS SOA model is computationally most expensive option. Unfortunately, we did not record the actual computational cost for each of the SOA schemes or the gas-phase chemistry schemes.

The end of section 1.2 Would it be possible for the authors to comment on what conditions the activation schemes are initialised? Running at higher than 1km, presumably the assumption is to use the aerosol composition, both SOA and SIA, at a specific RH which is then fed into the ARG or FN schemes with regards to hygroscopicity? This might also impact the performance of any given activation scheme if the assumed mass is from a 'dry' SOA partitioning model, whereas SIA accounts for RH dependent partitioning.

Reply: The chemical initial and boundary conditions (ICONs/BCONs) come from the modified CESM/CAM version 5.3 with updates by Gantt et al. (2014), He and Zhang (2014), and Glotfelty et al. (2016). Only the SIA concentrations are present as ICONs/BCONs. A 10-day spin-up is also used for the SOA concentrations to stabilize.

Section 3.1: How does the new treatment of semi-volatile POA work with boundary conditions used to initialise simulations? What is the impact of forcing different VBS profiles into one? I wasn't clear how

this relates to, for example, the inputs required for the CCN schemes. Is it related to an inability to track separate sources through the simulation? Or is it related to how the emissions are convolved?

Reply: As mentioned above, other than the differences in POA emissions as mentioned in the text, there are no other differences in ICONs/BCONs used in all the VBS cases. The model is unable to track the use of different VBS profiles in 1 simulation. Different simulations would have to be run, each changing 1 parameter to understand how the particular change affects SOA concentrations and CCN. The standard model inputs apply for both the VBS schemes, and the CCN schemes. No other special model inputs are required to run the CCN scheme, as the CCN scheme is dependent on the aerosol concentrations from the chemistry schemes, and vice versa.

Section 4.3 Related to a previous point, the authors comment on how larger differences in CDNC predictions arise from different gas-phase mechanisms over VBS variants. I think it would benefit the reader, and the context of the sensitivity simulations to comment on how the VBS versus RH interactions feed into the CDNC parametrisations.

Reply: The evaluation of the performance of RH by the model is interesting but beyond the scope of this study. However, from previous research, for example, by Yahya et al. (2015, 2016), the model performs relatively well for RH. We also do not expect RH to vary much between the different VBS cases. To ultimately understand how RH impacts the VBS performance, we would have to artificially vary the RH as inputs to the model. This might work better as a box-model study, rather than for a 3d model, where RH is predicted, and is temporally and spatially varying.

References: Yahya, K., J. He, and Y. Zhang (2015), Multiyear applications of WRF/Chem over continental U.S.: Model evaluation, variation trend, and impacts of boundary conditions, J. Geophys. Res. Atmos., 120, 12748–12777, doi:10.1002/2015JD023819.

Yahya, K., Wang, K., Campbell, P., Glotfelty, T., He, J., and Zhang, Y.: Decadal evaluation of regional climate, air quality and their interactions over the continental US using WRF/Chem version 3.6.1, Geosci. Model Dev., 9, 671 – 695, doi:10.5194/gmd-9-671-2016, 2016

A more philosophical question, which doesn't require any modifications and isn't a critique of this study: I often wonder how much value we should place on assuming accurate ambient OA/OC measurements without going back to trailing the same model permutations in a controlled environment. Would the authors value smog chamber studies, on mixed VOC systems, using the same parametrisations but in a box-model configuration? It seems that, at least, this would be valuable from a high accuracy measurements perspective.

Reply: Smog chamber studies, as well as box-model configurations are definitely valuable. In our opinion, they offer complementary information to 3-D model testing. As a matter of fact, many smog chamber and box-model studies were indeed carried out first, before the incorporation of the derived parameterisations into a 3-D model such as the WRF/Chem model in this study. While box model studies are confined to a controlled environment, using the 3-D model in our case represents real atmosphere

yet it introduces other uncertainties to OA concentrations from other atmospheric variables, feedbacks and processes. Both smog chamber studies/box-model studies and 3-D model studies have their own purposes and strengths, and should be used when resources are permitted.

Minor comments: Abstract line 41:'to 7.1%, it, however'. Please break the sentence here

Reply: This has been modified.

page 14 line 316-317:'based on a number of literature', should be 'based on a number of studies in the literature'

Reply: This has been modified.

page 29, lines 661-662:'simulated vs, observed' please replace this with 'simulated versus observed'

Reply: This has been modified.

1 2 3	Modeling Regional Air Quality and Climate: Improving Organic Aerosol and Aerosol Activation Processes in WRF/Chem version 3.7.1
4	Khairunnisa Yahya <sup>1</sup> , Timothy Glotfelty <sup>1</sup> , Kai Wang <sup>1</sup> , Yang Zhang <sup>1</sup> *, and Athanasios Nenes <sup>2,3,4,5</sup>
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15 16	Email: *yang_zhang@ncsu.edu
17	ABSTRACT
18	Air quality and climate influence each other through the uncertain processes of aerosol formation

19 and cloud droplet activation. In this study, both processes are improved in the Weather, Research 20 and Forecasting model with Chemistry (WRF/Chem) version 3.7.1. The existing Volatility Basis 21 Set (VBS) treatments for organic aerosol (OA) formation in WRF/Chem is improved by 22 considering the secondary OA (SOA) formation from semi-volatile primary organic aerosol 23 (POA), a semi-empirical formulation for the enthalpy of vaporization of SOA, as well as 24 functionalization and fragmentation reactions for multiple generations of products from the 25 oxidation of VOCs. Two-month long simulations (May to June 2010) are conducted over 26 continental U.S. and results are evaluated against surface and aircraft observations during the 27 Nexus of Air Quality and Climate Change (CalNex) campaign. Among all the configurations 28 considered, the best performance is found for the simulation with the 2005 Carbon Bond 29 mechanism (CB05) and the VBS SOA module with semivolatile POA treatment, 25%

30 fragmentation, and the emissions of semi-volatile and intermediate volatile organic compounds 31 being 3 times of the original POA emissions. Among the three gas-phase mechanisms (CB05, 32 CB6, and SAPRC07) used, CB05 gives the best performance for surface ozone and  $PM_{2.5}$ 33 concentrations. Differences in SOA predictions are larger for the simulations with different VBS 34 treatments (e.g., non-volatile POA <del>vs.</del>versus semivolatile POA) as compared to the simulations 35 with different gas-phase mechanisms. Compared to the simulation with CB05 and the default SOA 36 module, the simulations with the VBS treatment improve cloud droplet number concentration 37 (CDNC) predictions (NMBs from -40.8% to a range of -34.6% to -27.7%), with large differences 38 between CB05/CB6 and SAPRC07 due to large differences in their OH and HO<sub>2</sub> predictions. An 39 advanced aerosol activation parameterization based on the FN05 series reduces the large negative 40 CDNC bias associated with the default ARG00 parameterization from -35.4% to a range of -0.8% 41 to 7.1%. It, however, increases the errors due to overpredictions of CDNC, mainly over 42 northeastern U.S. This work indicates a need to improve other aerosol-cloud-radiation processes 43 in the model such as the spatial distribution of aerosol optical depth and cloud condensation nuclei 44 in order to further improve CDNC predictions.

# 45 **1. Introduction**

46

The Intergovernmental Panel on Climate Change (IPCC) report on the AR5 scenario attributes the aerosol radiative forcing (RF) to be the dominant source of uncertainty contributing to the overall uncertainty in the net Industrial Era Radiative Forcing (RF) calculations (Myhre et al., 2013). Despite the inclusion of more aerosol processes in the current generation of atmospheric models, differences between atmospheric models and observations continue to persist. Aerosols affect the climate through the direct effect by absorbing or scattering radiation, or the indirect effect by acting as cloud condensation nuclei (CCN). According to Hallquist et al. (2009), the 54 formation of inorganic particulates such as sulfate, nitrate, and ammonium are well understood, 55 however, there are large uncertainties in the formation of secondary organic aerosol (SOA). As a 56 result, current models do not have a comprehensive treatment of OA, which usually result in an 57 underprediction of OA concentrations (Hodzic et al., 2010; Jathar et al., 2011; Bergstrom et al., 58 2012), due to missing key precursors and processes in OA formation (Ahmadov et al., 2012). Some 59 of the missing key precursors in most models include semi-volatile primary organic aerosol (POA), 60 long-chain *n*-alkanes, polycyclic aromatic hydrocarbons (PAHs), and large olefins that have lower 61 volatilities compared to traditional SOA precursors (Chan et al., 2009). The organic carbon (OC) 62 component of the radiative forcing in the IPCC AR5 report also does not include SOA with the 63 reason that the formation is dependent on a number of factors that are not currently sufficiently 64 quantified (Myhre et al., 2013). However, SOA can form a significant percentage of total OA (up 65 to 95% in rural areas) (Zhang et al., 2007). Another large source of uncertainty is the quantification of clouds as well as aerosol-cloud interactions due to incomplete or inaccurate representations of 66 67 these processes in climate models (Boucher et al., 2013). A major process in cloud formation from 68 aerosol is aerosol activation, which involves the condensational growth of aerosols in a cooling air 69 parcel until maximum supersaturation, and some of the wet particles reach a critical radius where 70 they are then able to grow spontaneously into cloud droplets (Ghan et al., 2011). Various 71 approaches have been developed to reduce the uncertainties associated with OA and aerosol 72 activation treatments in climate models. Those treatments are reviewed in the following section.

73 74

# 1.1. VBS Treatments and Sensitivity to Different Gas-Phase Chemical Mechanisms in Regional and Global Models

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Unlike inorganic aerosols such as sulfate, the physical and chemical properties of OA
dynamically evolve with age (Jimenez et al., 2009). The traditional approach to modeling SOA is

78 to assume that each VOC precursor forms several surrogate compounds (Odum et al., 1996). 79 However, the traditional method has several shortcomings, for example, two products are needed 80 for each VOC precursor causing this method to be computationally-expensive if many VOC 81 precursors are treated in the model (Murphy and Pandis, 2009). The assumption that the products 82 are unreactive also does not reflect the dynamic nature of the first generation products from the 83 oxidation of VOCs that can undergo successive oxidation steps to further produce lower volatility 84 products (Jimenez et al., 2009). The volatility basis set (VBS) is a framework developed by 85 Donahue et al. (2006), which is able to simulate gas-phase partitioning and multiple generations of gas-phase oxidation of organic vapors. This approach addresses the shortcomings of the 86 87 traditional SOA modeling approach as it can cover the complete volatility range of OA compounds 88 (Murphy and Pandis, 2009).

89 Table 1 summarizes some of the VBS treatments from current regional and global models. The 90 VBS treatment has been implemented into a number of regional models such as the Weather, 91 Research and Forecasting model with Chemistry (WRF/Chem) (Shrivastava et al., 2011; 92 Ahmadov et al., 2012), the Particulate Matter Comprehensive Air Quality Model with extensions 93 (PMCAMx) (Lane et al., 2008; Donahue et al., 2009; Murphy et al., 2009), and CHIMERE (Hodzic 94 et al., 2010). It has also been implemented in global models such as GISS II' GCM (Farina et al., 95 2010; Jathar et al., 2011) and the Community Earth System Model (CESM) (Shrivastava et al., 96 2015). Different studies define the classifications of the organic species slightly differently. Donahue et al. (2009) defined primary organic vapors with effective saturation concentrations (C\*) 97 of  $10^{-2}$  -  $10^{-1}$ ,  $10^{0}$  -  $10^{2}$ , and  $10^{3}$  -  $10^{6}$  µg m<sup>-3</sup>at 298 K to be low volatility organic compounds 98 99 (LVOCs), semi-volatile organic compounds (SVOCs), and intermediate volatility organic 100 compounds (IVOCs), respectively. Shrivastava et al., (2011) and Jathar et al. (2011) defined

101 primary organic vapors with C\* values of  $10^{-2} - 10^3$  and  $10^4 - 10^6 \,\mu g \,m^{-3}$  to be SVOCs and IVOCs, 102 respectively. All those studies defined VOCs to be gas-phase organic species with C\* larger than 103  $10^6 \,\mu g \,m^{-3}$  at 298 K.

104 The traditional emission inventories used in the chemical transport models consist of VOCs 105 but not SVOCs or IVOCs as both SVOCs and IVOCs are difficult to measure. This is most likely 106 because SVOCs and IVOCs tend to evaporate at high temperatures from combustion sources 107 (Donahue et al., 2009). As the traditional SOA approach usually underpredicts the SOA 108 concentration, the addition of the SVOC and IVOC emissions on top of the existing VOC 109 emissions in most emission inventories can improve model performance. To account for the 110 missing key precursors in OA formation, SVOC and IVOC emissions are usually estimated as a 111 factor of existing POA emissions in current emission inventories. For example, Shrivastava et al. 112 (2011) estimated the sum of all SVOC and IVOC precursors to be 7.5 times the mass of traditional 113 POA emissions inventory over Mexico City, but indicated that the scaling factor of 3 for SVOC 114 emissions based on the POA emissions is poorly constrained. Shrivastava et al. (2008) and Jathar 115 et al. (2011) assumed that SVOC emissions are represented by the traditional emission inventory 116 while IVOC emissions are 1.5 times the traditional emission inventory. Pye and Seinfeld (2010) 117 assumed that SVOC emissions are a subset of traditional POA emission inventories, and their POA 118 emissions were scaled up by 27% on a global scale. IVOC emissions are assumed to be spatially-119 distributed similar to naphthalene and are predicted to be roughly a factor of half of global POA 120 emissions. Tsimpidi et al. (2014) assumed that the IVOC emissions are 1.5 times the traditional POA emission inventory and are assigned to the 4<sup>th</sup> volatility bin with  $C^* = 10^5 \square g m^{-3}$ . For 121 122 comparison, some studies such as Ahmadov et al. (2012) and Bergstrom et al. (2012) used the 123 VBS approach for OA modeling but did not include additional SVOC emissions. There are also

differences in the volatility distribution used in literature. Shrivastava et al. (2008) and Jathar et al. (2011) found that moving half the mass of SVOC from all bins to the lowest bin from the traditional "diesel exhaust" volatility distribution of Robinson et al. (2007) produced the lowest errors in simulated OA on an annual average basis.

The number of bins used can also result in differences in simulated SOA concentrations. Shrivastava et al. (2011) showed that the 2-species VBS performed better than the 9-species VBS in modeling oxygenated organic aerosol (OOA) and gave the closest agreement to the OOA calculated by the Positive Matrix Factorization (PMF) method. This indicates that SOA may be less volatile as compared to the volatility distribution in the 9-species VBS which allows for evaporation of SOA with dilution (Shrivastava et al., 2011).

134 The amount of oxygen added for each oxidation step may contain uncertainties. This factor 135 can influence the O:C ratio used for the model evaluation. O:C predictions from models need to 136 be improved by including fragmentation reactions (which could lead to an increase in O:C ratios) 137 and improving emission estimates (Shrivastava et al., 2011). Different rate constants can also result 138 in different predictions of SOA concentrations. For example, Farina et al. (2010) showed that the use of k value of  $1 \times 10^{-12}$  cm<sup>3</sup> molecule<sup>-1</sup> s<sup>-1</sup> compared to the default k value of  $10 \times 10^{-12}$  cm<sup>3</sup> 139 molecule<sup>-1</sup> s<sup>-1</sup> resulted in a reduced aged SOA formation by 71%. Hodzic et al. (2010) also showed 140 141 a case study based on Grieshop et al. (2009) in which each oxidation step reduced the volatility of 142 the S/IVOC vapors by two orders of magnitude and each successive oxidation step produced a 143 40% increase in mass due to the addition of oxygen. This case is inconclusive in urban areas - a 144 larger bias along with a higher correlation coefficient compared to the more common case where 145 each oxidation step reduced the volatility by one order of magnitude with a 7.5% increase in mass.

However, the model performed worse (with larger bias and lower correlation coefficient) insuburban areas.

The aging process improves model performance in general in the United States (U.S.) but deteriorates the performance in several parts of Europe. Accounting for the aging process of OA will increase the OA concentrations and improve model results in the U.S. where OA is usually underpredicted, but increase the model bias for OA in several parts of Europe where OA concentrations are overpredicted (Farina et al., 2010; Bergstrom et al., 2012).

153 Shrivastava et al. (2013) studied the effects of the fragmentation and functionalization in VBS. 154 Functionalization increases the mass of OA for each successive oxidation step, while 155 fragmentation reduces the mass for each oxidation step. One such a case includes simulating first-156 order effects of the fragmentation and functionalization processes in VBS by assuming 157 functionalization of 100% of organic vapors for the first two generations of oxidation and both 158 fragmentation and functionalization for the third and higher generations of oxidation. The 159 fragmentation reduces the SOA concentrations drastically. For example, Shrivastava et al. (2013) 160 showed that peak SOA concentrations can be reduced by factors of 2 to 4 for a 1-hour example on 161 10 March 2006 at 21 UTC over Mexico City Plateau.

The VBS framework for OA modeling in the latest version of WRF/Chem, v3.7.1, is coupled with several gas-phase mechanisms including the 2005 Carbon Bond Mechanism (CB05) (Yarwood et al., 2005), the Model for Ozone and Related chemical Tracers version 4 (MOZART-4) (Emmons et al., 2010), the Regional Atmospheric Chemistry Model (RACM) (Stockwell et al., 166 1997), and the 1999 version of the Statewide Air Pollution Research Centre (SAPRC99) 167 mechanism (Carter, 2000). Different gas-phase mechanisms have different lumpings/groupings 168 for VOCs, which will affect OA formation. For example, VOCs are lumped according to their

169 carbon bonds (e.g., single or double bond) in CB05 (Yarwood et al., 2005) while VOCs in 170 SAPRC99 (Carter, 2000) are lumped according to their OH reactivities. A number of studies 171 examined the differences in predicting O<sub>3</sub> concentrations due to different gas-phase mechanisms 172 (e.g., Luecken et al., 2008; Li et al., 2012; Shearer et al., 2012; Zhang et al., 2012), but fewer 173 studies reported the impact of different gas-phase mechanisms on modeling SOA and PM<sub>2.5</sub> 174 concentrations (Kim et al., 2011; Zhang et al., 2012). SAPRC99 has more detailed organic 175 chemistry compared to CB05. SAPRC99 has been updated to SAPRC07 (and recently, to 176 SAPRC11) based on newly available information regarding the reactions and influence of 177 individual VOCs on  $O_3$ , as well as evaluations against chamber experiments (Carter, 2010). In 178 addition, SAPRC07 has reformulated reactions of peroxy radicals so that the effects of changes in 179 nitrogen oxides (NO<sub>x</sub>) on organic product formation is more accurately represented. SAPRC07 180 has the most extensive set of VOC species and reactions, as compared to CB05 and the Carbon 181 Bond version 6 (CB6). Shearer et al. (2012) reported that a condensed version of SAPRC07 182 predicted lower O<sub>3</sub> and OH concentrations in central California compared to SAPRC99 due to a 183 decreased reaction rate coefficient in the reaction of OH and NO<sub>2</sub> to form HNO<sub>3</sub>. Li et al. (2012) 184 also showed that predicted O<sub>3</sub> concentrations from SAPRC07 were lower than those of SAPRC99 185 by up to 20% over Texas. The same study also reported that SAPRC07 gave lower OH 186 concentrations due to differences in the reaction rate constants in the reactions of  $O^{1}D$  and  $H_{2}O$ 187 between SAPRC07 and SAPRC99. Luecken et al. (2008) reported that SAPRC99 gave higher O<sub>3</sub> 188 concentrations compared to CB05 on average; however, the differences vary with locations, 189  $VOC/NO_x$  ratios, and the concentrations of precursor pollutants. This is consistent with the results 190 from Zhang et al. (2012), which predicted that SAPRC99 using WRF/Chem with the Model of 191 Aerosol, Dynamics, Reaction, Ionization and Dissolution (WRF/Chem-MADRID) produced the

192 highest O<sub>3</sub> mixing ratios in July at the Southeastern Aerosol Research and Characterization 193 (SEARCH) sites. The CB6 (Yarwood et al., 2010) is an updated version of CB05 with improved 194 kinetic and photolysis data, additional explicit species for long-lived and abundant organic 195 compounds including propane, acetone, benzene and acetylene, as well as revised isoprene and 196 aromatics chemistry from CB05. Yarwood et al. (2010) showed that CB6 produces higher daily 197 maximum 8-hr O<sub>3</sub> as compared to CB05 over Los Angeles for one episode day in August with the 198 highest observed O<sub>3</sub> mixing ratios. CB6 was also shown to produce substantially higher OH concentrations (25% to 50% higher at mid-day over large areas) over eastern U.S. compared to 199 200 CB05 over a few days in June, 2006. A summary of the main characteristics of CB05, CB6, and 201 SAPRC07 gas-phase mechanisms are listed in Table 2.

### 202 **1.2. Description of Aerosol Activation Parameterizations**

203

204 Ghan et al. (2011) provided a comprehensive review on various aerosol activation treatments 205 in current climate models. Two main types of parameterizations are commonly used: the Abdul-206 Razzak and Ghan (2000) (AR-G00) and the Fountoukis and Nenes (2005) (FN05) and associated 207 updates described in Barahona et al. (2010) and Morales Betancourt and Nenes (2014). AR-G00 208 uses multiple lognormal or sectional distributions to approximate the aerosol size distribution. It 209 uses the Kohler theory to relate the aerosol size distribution and composition to the number of 210 aerosols activated as a function of maximum supersaturation  $(S_{max})$ . The complex function 211 involving  $S_{max}$  is parameterized based on the standard deviation  $\sigma$  from a large number of numerical solutions using a cloud parcel model. The number and mass activated are particles with 212 213 critical supersaturation less than  $S_{max}$ . It also accounts for particle growth before and after the 214 particles are activated. However, the ARG treatment does not explicitly represent kinetic 215 limitations which tend to affect smaller or larger particles (with diameters far from their critical

216 size). Very small particles tend to lose water when supersaturation declines as they never exceed 217 the critical supersaturation for that particle size, and very large particles may not have achieved 218 the critical size before  $S_{max}$  is reached (Ghan et al., 2011). Kinetic limitations refer to the (i) inertial 219 mechanism – where particles with large dry diameters grow to be as large as activated particles 220 but have not been activated themselves, these particles should be considered together with 221 activated particles; (ii) evaporation mechanism – where particles with high critical supersaturation 222 evaporate before reaching their critical diameters; and the (iii) deactivation mechanism – where 223 initially activated particles that are deactivated to interstitial aerosols when the parcel 224 supersaturation falls below the equilibrium supersaturation (Nenes et al., 2001). Neglecting kinetic 225 limitations performs well for all conditions except in highly-polluted areas (Ghan et al., 2011). In 226 urban and highly-polluted cases, many particles fail to be activated due to strong evaporation and 227 deactivation processes (Nenes et al., 2001). Explicitly accounting for kinetic limitations reduces 228 CDNC at low updraft velocity (Ghan et al., 2011).

229 The Fountoukis and Nenes (2005) (FN05) scheme improved the ARG00 scheme by solving 230 S<sub>max</sub> analytically (with the exception of kinetically-limited particles) using a so-called "population 231 splitting" method. In addition, FN05 took into account the kinetic limitations, as well as the 232 influence of gas kinetics on water vapor diffusivity (Ghan et al., 2011). The other improved 233 treatments built on top of the FN05 scheme include the entrainment of ambient air, which could 234 reduce the supersaturation of the updraft (Barahona and Nenes, 2007) (BN07) (therefore reducing 235 CDNC); the adsorption of water vapor onto insoluble particles by Kumar et al. (2009) (KU09) 236 based on a modified Frenkel-Halsey-Hill (FHH) adsorption theorem (which will increase CDNC); 237 the growth of giant cloud condensation nuclei (CCN) (Barahona et al., 2010) (BA10) by 238 introducing an additional condensation rate term to account for condensation of giant CCN (which will reduce CDNC); as well as the modification of the original population splitting concept in
FN05 and BA10 by Morales Betancourt and Nenes (2014) (MN14) by better accounting for the
size of inertially limited CCN, and removing a discontinuity in the calculation of the surface area
of cloud droplets.

The parameterization of Abdul Razzak and Ghan (2000) (ARG00) is used as the default aerosol activation module in WRF/Chem. It is not linked to the microphysics module or cumulus parameterization in WRF or WRF/Chem. However, for WRF/Chem, the cloud droplet number concentration (CDNC) generated in ARG00 is passed to the microphysics scheme, i.e., the Morrison two-moment microphysics scheme selected in this work.

248 249

## **1.3 Motivations and Objectives**

250 The online-coupled meteorology and chemistry model, WRF/Chem, has recently been 251 applied for air quality and climate modeling on a decadal scale (Yahya et al., 2016a, b). 252 WRF/Chem can also simulate aerosol direct and indirect feedbacks, which are important 253 considerations for climate modeling. However, as mentioned previously, the representations of 254 OA and aerosol-cloud interactions in current regional and global climate models are subject to 255 large uncertainties. In particular, while the VBS framework in WRF/Chem significantly improves 256 SOA performance (Wang et al., 2015), it lacks the semi-volatile POA treatment, as well as 257 fragmentation processes (Shrivastava et al., 2013). The first objective of this study is to reduce 258 uncertainties associated with OA predictions by improving the existing VBS module in 259 WRF/Chem and identifying the best gas-phase chemical mechanism to drive the VBS module for 260 the most accurate OA predictions. The impact of the improved OA predictions on CDNC in 261 WRF/Chem will be quantified. The second objective is to incorporate an improved aerosol activation parameterization based on the FN05 series into WRF/Chem to study its impacts onCDNC predictions.

### 264 2. Model Configuration, Evaluation Protocol, and Observational Datasets

265

# 266 2.1.Model Setup and Inputs267

268 The model used in this study is a modified version of WRF/Chem v3.7.1 as described by Wang 269 et al. (2015). The 2005 Carbon Bond gas-phase mechanism (CB05) of Yarwood et al. (2005) with 270 additional chlorine chemistry is coupled with the Modal for Aerosol Dynamics in Europe – 271 Secondary Organic Aerosol Model (MADE/SORGAM) (Ackermann et al., 1998; Schell et al., 272 2001) and the Volatility Basis Set (MADE/VBS) (Ahmadov et al., 2012). The CB05-VBS option 273 has also been coupled to existing model treatments including the aerosol direct effect, the aerosol 274 semi-direct effect on photolysis rates of major gases, and the aerosol indirect effect on CDNC and 275 resulting impacts on shortwave radiation. The physics options used in WRF/Chem include the 276 rapid and accurate radiative transfer model for GCM (RRTMG) for both shortwave and longwave 277 radiation, the Yonsei University (YSU) planetary boundary layer (PBL) scheme (Hong et al., 2006; 278 Hong, 2010), the Morrison et al. (2009) double moment microphysics scheme, as well as the Multi-279 scale Kain-Fritsch (MSKF) cumulus parameterization scheme (Zheng et al., 2016). Aqueous-280 phase chemistry module (AQCHEM) for both resolved and convective clouds is based on a similar 281 AQCHEM module in CMAQv4.7 of Sarwar et al. (2011). The anthropogenic emissions used are 282 from the 2010 emissions based on the 2008 U.S. Environmental Protection Agency (U.S. EPA) National Emissions Inventory (NEI) from the Air Quality Model Evaluation International Initiative 283 284 (AQMEII) project (Pouliot et al., 2015). Dust emissions are based on the Atmospheric and 285 Environmental Research Inc. and Air Force Weather Agency (AER/AFWA) scheme (Jones and 286 Creighton, 2011). Emissions from sea salt are generated based on the scheme of Gong et al. (1997).

Biogenic emissions are simulated online by the Model of Emissions of Gases and Aerosols from
Nature v2.1 (MEGAN2.1) (Guenther et al., 2006).

289 The chemical initial and boundary conditions (ICONs/BCONs) come from the modified 290 CESM/CAM version 5.3 with updates by Gantt et al. (2014), He and Zhang (2014), and Glotfelty 291 et al. (2016). The meteorological ICONs/BCONs are from the National Center for Environmental 292 Protection Final Reanalyses (NCEP FNL) dataset, which is available every 6 hours. The chemical 293 fields are also allowed to run continuously while the meteorology is reinitialized every 5 days. The 294 simulations are performed at a horizontal resolution of 36-km with  $148 \times 112$  horizontal grid cells 295 over the CONUS domain and parts of Canada and Mexico, and a vertical resolution of 34 layers 296 from the surface to 100-hPa.

A number of sensitivity simulations are designed to identify the model configuration with results that are in the closest agreement to observations as well as the realistic model treatments of OA that are the closest to atmospheric processes. The baseline and sensitivity simulations are conducted from May to June 2010, during which the Nexus of Air Quality and Climate Change (CalNex) campaign was held in Bakersfield and Pasadena, California. The first 10 days from May 1<sup>st</sup> to May 10<sup>th</sup> are considered to be the spin-up period.

303 304

# 2.2. Model Evaluation Protocol and Available Measurements

305 Statistical measures including the Mean Bias (MB), Correlation Coefficient (Corr), 306 Normalized Mean Bias (NMB) and Normalized Mean Error (NME) (Yu et al., 2006) are used to 307 evaluate the simulations against observational data. Observational data are available for organic 308 carbon (OC) and total carbon (TC) from the Speciated Trends Network (STN) and the Interagency 309 Monitoring for Protected Visual Environments (IMPROVE). While both OC and TC from 310 IMPROVE are used for model evaluation, only TC data from STN are used as STN uses the 311 thermo-optical transmittance protocol for OC that is different from the one used by IMPROVE 312 (Zhang et al., 2012). In addition, the measurements for STN OC are not blank corrected for carbon 313 on the background filter (Wang et al., 2012). The ratios OA/OC ratios vary across locations in the 314 continental U.S. (CONUS) depending on whether the OA is dominated by secondary formation 315 (Aitken et al., 2008) or it contains more aliphatic hydrocarbons (Turpin and Lim, 2001). In this 316 study, two ratios, 1.4 and 2.1, are used to convert simulated OA to OC based on a number of 317 studies in literature (Turpin and Lim, 2001; Aitken et al., 2008; Xu et al., 2015). As the simulations 318 are based on CONUS with varying OA properties (less or more oxidized OA), the use of two 319 OA/OC ratios can represent the different types of OA present for all locations in the U.S. Spatial 320 plots, time series plots at specific sites, as well as overlay plots are used to evaluate model 321 performance. The IMPROVE sites chosen for the time series plots include the visibility-protected 322 areas in Brigantine National Wildlife Refuge (NWR), NJ, Death Valley National Park (NP), CA, 323 Swanqwarter National Wildlife Refuge (NWR), NC, and the Tallgrass Prairie National Preserve, 324 KS. The Brigantine NWR is a tidal wetland and has a shallow bay, the Death Valley NP is a desert, 325 and the Swangwarter NWR is a coastal brackish marsh. The time series plots are made at four STN 326 sites including two urban sites: in Washington, DC and Boise, ID, one industrial site in Tampa, 327 FL, and one rural/agricultural site in Liberty, KS. SOA, hydroxyl radical (OH) and hydroperoxy 328 radical (HO<sub>2</sub>) data are also available for May to June 2010 as part of the California Research at the 329 Nexus of Air Quality and Climate Change (CalNex) campaign (Kleindienst et al., 2012; 330 Lewandowski et al., 2013) in Bakersfield, CA and Pasadena, CA, which are both urban locations. 331 The Bakersfield sampling site is located between the city center and areas of agricultural activity, 332 while the Pasadena site is located at the California Institute of Technology campus within the Los 333 Angeles metropolitan area to the southwest and mountains in the north (Baker et al., 2015).

334 POA/OA ratios are also used to evaluate the performance of the model. A number of studies 335 have reported observed POA/OA ratios which range from 15% to 40% over CONUS. For 336 example, over southeastern U.S., hydrocarbon-like OA (HOA) and cooking OA are found to 337 contribute to 21 - 38% of total OA in urban sites (Xu et al., 2015). HOA and oxygenated OA 338 (OOA) are found to account for 34% and 66% of measured OA from Pittsburgh in September 2002 339 (Zhang et al., 2005). HOA and cooking OA are assumed to be synonymous to POA, and OOA is 340 assumed to be synonymous to SOA. Particulate matter sampled during August and September 341 2006 in Houston as part of the Texas Air Quality Study II Radical and Aerosol Measurement 342 Project showed that approximately 32% of OA comes from HOA (Cleveland et al., 2012). Results 343 from positive matrix factorization analysis from the Pasadena ground site during May and June 344 2010 showed that the primary components contribute 29% of the total OA mass (Hayes et al., 345 2013). Based on Zhang et al. (2007), the percentages of HOA mass at urban sites in Riverside, 346 CA, from mid-July to mid-August 2005, in Houston, TX, from mid-August to mid-September 347 2000, and in New York City in July 2001 are 15%, 38%, and 30%, respectively. In addition, Zhang 348 et al. (2011) compiled a large number of field campaigns across the globe where the average 349 POA/OA ratios for urban, downwind and rural/remote areas are found to be 0.42, 0.18 and 0.10 350 respectively.

For the aerosol activation sensitivity and production simulations, additional variables that will be analyzed in this study include maximum 1-hour and 8-hour O<sub>3</sub> against the Clean Air Status and Trends Network (CASTNET) and Air Quality System (AQS), aerosol optical depth (AOD), CDNC and cloud condensation nuclei (CCN) against MODIS.

355 3. Model Development and Improvement356

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357 A number of modifications have been made to the standard version of WRF/Chem model 358 v3.7.1. Those modifications and treatments are described below.

- 359 **3.1. OA Treatments**
- 360

361 The CB05-VBS treatment in the default WRF/Chem v3.7.1 assumes that POA is nonreactive 362 and nonvolatile. In this study, POA is assumed to be semivolatile, and can undergo gas-particle 363 partitioning, similar to anthropogenic SOA (ASOA) and biogenic SOA (BSOA) in VBS. While the volatility of ASOA and BSOA is represented by 4 bins with C\* from  $10^0$  to  $10^3$  µg m<sup>-3</sup>. The 364 POA is distributed into 9 bins, with C\* from  $10^{-2}$  to  $10^{6}$  µg m<sup>-3</sup>, following the set-up of Shrivastava 365 366 et al. (2011). The POA is oxidized to form semi-volatile OA (SVOA), which can also undergo 367 gas-particle partitioning. For the POA, bin-resolved enthalpies of vaporizations are used, ranging from 64 kJ mol<sup>-1</sup> for the 9<sup>th</sup> bin to 112 kJ mol<sup>-1</sup> for the 1<sup>st</sup> bin according to Shrivastava et al. (2011). 368 The default enthalpy of vaporization ( $\Delta H_{vap}$ ) for SOA in WRF/Chem is 30 kJ mol<sup>-1</sup> according to 369 Lane et al. (2008). A more accurate alternative is to use the  $\Delta H_{vap}$  values calculated from the semi-370 371 empirical correlation from Epstein et al. (2010):

$$\Delta H_{\rm vap} = -11 \, \log_{10} \, \mathrm{C}^*_{300} + 129 \tag{1}$$

373 The values of  $\Delta H_{vap}$  Epstein et al. (2010) are used in a number of sensitivity simulations and final 374 production simulation.

375 Shrivastava et al. (2013, 2015) also implemented several cases of fragmentation and 376 functionalization (FF) processes into VBS. For this study, the FF set-up is similar to the method 377 employed by Shrivastava et al. (2013), with the exception that fragmentation percentages of 10%, 378 25%, and 50% are used in sensitivity simulations. Shrivastava et al. (2013) used fragmentation 379 percentages of 50% (intermediate fragmentation) and 85% (high fragmentation) in his simulations 380 over Mexico City. For example, for the 10% FF case, 10% of the mass in the VBS species is functionalized and moved to the next lower volatility bin, 80% is fragmented and moved to the highest volatility bin, and the remaining 10% is fragmented and becomes more volatile than the highest volatility bin (i.e., it is lost from the current volatility bins). For the 50% FF case, 50% is functionalized and moved to the next lower volatility bin, 40% is fragmented and moved to the highest volatility bin, and 10% is lost.

2hao et al. (2014) measured IVOCs in Pasadena, CA during CalNex and found that the concentrations of primary IVOCs are similar to those of single-ring aromatics, and they produce about 30% of newly formed SOA in the afternoon. With the semivolatile POA and FF cases in this study, additional IVOC and SVOC emissions are added as three times of the traditional POA emissions from NEI, to account for missing IVOC and SVOC species in the traditional POA emission inventory. The fraction of IVOC/SVOC emissions assigned to each volatility bin is summarized in Table 3.

393 The mass fraction of organics in each volatility bin determined in laboratory studies also differs significantly according to the sources of organics. For example, May et al. (2013a, b, c) has 394 395 different volatility distributions of mass fractions of organics for gasoline vehicle exhaust, diesel 396 exhaust, and biomass burning. To take into account the different sources of organic compounds 397 into a single volatility distribution for the purpose of this work, a new volatility distribution is 398 calculated based on the mass fractions reported by Shrivastava et al. (2011), May et al. (2013a, c) 399 and the percentages of VOC emissions from various sources from the 2008 NEI. According to the 400 2008 NEI report (Rao et al., 2013), total VOC emissions from stationary, mobile and fire 401 (prescribed and wildfire) sources are ~7.6, ~5.6, and ~49.6 million tons, respectively. The 402 corresponding percentages for VOC emissions are ~12%, ~9%, and ~79% for stationary, mobile, 403 and fire sources, respectively. Based on the U.S. EPA (2013), the percentages of diesel emissions

404 from mobile sources are low compared to gasoline sources (~7% of total diesel and gasoline
405 sources); they are thus not included in this study.

406 An example calculation for the mass fraction of the lowest volatility bin for POA and
407 IVOC/SVOC emissions are as follows:

408 
$$\log C_{-2}^{*}$$
 (at 298K) = 0.04×12% + 0.14×9%+0.79×79% = 0.1754 (2)

where  $C_{-2}^*$  refers to the lowest volatility bin with a value of  $10^{-2} \mu g m^{-3}$ , 12%, 9%, and 79% refer 409 410 to the percentages for VOC emissions from stationary, mobile, and fire sources, respectively from 411 NEI, 0.04 refers to the original mass fraction for stationary emissions based on anthropogenic emissions from Shrivastava et al. (2011) for the lowest volatility bin with a value of  $10^{-2} \mu g m^{-3}$ , 412 413 0.14 refers to the original mass fraction for gasoline emissions from May et al. (2013a) for the lowest volatility bin with a value of  $10^{-2} \,\mu g \, m^{-3}$ , 0.2 refers to the original mass fraction for biomass 414 burning emissions from May et al. (2013c) for the lowest volatility bin with a value of  $10^{-2} \mu g m^{-1}$ 415 416 <sup>3</sup>, and 0.1754 refers to the newly-calculated mass fraction of POA and IVOC/SVOC emissions for 417 this study. The mass fractions used by Shrivastava et al. (2011), May et al. (2013a, c), and this 418 work can be found in Table 3.

419 420

### **3.2.** Gas-Phase Chemical Mechanisms

Three gas-phase mechanisms are used: CB05, CB6, and SAPRC07. The gas-phase mechanisms for CB6 and SAPRC07 are coupled to the MADE/VBS in WRF/Chem v3.7.1 in this work following the coupling of CB05 with MADE/VBS by Wang et al. (2014). The emissions for all cases are based on the CB05 chemical species from the 2010 emissions based on the 2008 NEI. For SAPRC07, slight modifications had to be made to account for the different VOC species or groups. The mapping of emission species from CB05 to SAPRC07 is based on the grouping of 427 species from emitdb.xls from Henderson al. (2014)well from et as as 428 http://www.cert.ucr.edu/~carter/emitdb/old-emitdb.htm. CB05 emissions are used for the CB6 429 case, with the exception of the VOCs including propane, benzene, ethyne, acetone, and ketone that 430 are mapped based on fractions of existing CB05 VOCs according to Yarwood et al. (2010).

431 In VBS, the SOA precursors for CB6 are similar to those for CB05. The SOA precursors for 432 CB05 (and therefore CB6) are mapped from the default SAPRC99 precursors by Wang et al. 433 (2015). The SAPRC07 SOA precursors follow the existing mapping of SAPRC99-MOSAIC/VBS 434 in WRF/Chem. The chemical equations and rate parameters from ENVIRON (2013) and Carter 435 (2010) for CB6 and SAPRC07 gas-phase mechanisms, respectively, were included in the 436 chem/KPP/mechanisms directory in WRF/Chem. The SAPRC07 gas-phase mechanism 437 implemented in WRF/Chem in this case is the uncondensed and expanded version C, which 438 includes reactions for peroxy radical operators (Carter, 2010). Species in both CB6-MADE/VBS 439 and SAPRC07-MADE/VBS undergo dry deposition, aqueous chemistry, photolysis, and wet 440 scavenging that are similar to CB05-MADE/VBS.

- 441 **3.3.** Aerosol Activation
- 442

443 The FN05 series aerosol activation parameterizations (with the exclusion of MN14) have been 444 incorporated into 3-D regional air quality models and global climate and Earth system models such 445 as the WRF-Community Atmosphere Model version 5 (WRF-CAM5) (Zhang et al., 2015), and in 446 the global-through-urban WRF/Chem (GU-WRF/Chem) (Zhang et al., 2012) and CESM (Gantt et 447 al., 2014). In this study, the FN series parameterizations are incorporated into WRF/Chem 448 following the methods of Gantt et al. (2014) and Zhang et al. (2015) as described in detail in Zhang 449 et al. (2015). However, in WRF/Chem, the aerosol activation module is only linked to the microphysics module through the variable CDNC, which is read by the microphysics module. It 450

451 is not coupled to the cumulus parameterization scheme unlike in WRF-CAM5 and CESM. The 452 FN05 series has been incorporated into module\_mixactivate.F in the physics directory in 453 WRF/Chem. As BN07 involves the entrainment effect for convective clouds and has very small 454 impacts on non-convective CDNC (Zhang et al., 2015), it is not included in this study. In addition, 455 unlike Gantt et al. (2014) and Zhang et al. (2015), the KU09 treatment is also not included in this 456 study as the empirical constants A<sub>FHH</sub> and B<sub>FHH</sub> used in the formulation, which are compound-457 specific, have not been experimentally determined for black carbon, although those constants have 458 been determined for dust and confirmed by Laaksonen et al. (2016). The additional MN14 459 treatment incorporated in this study involves a small modification to the original FN05 series 460 parameterizations (without KU09), and helps to better account for the size of inertially limited 461 CCN, and to remove a discontinuity in the calculation of the surface area of cloud droplets 462 (Morales Betancourt and Nenes, 2014). The updated treatments are about 20% more 463 computationally expensive to run as compared to ARG00 (Zhang et al., 2016), but capture the 464 sensitivity of CDNC to all aspects of the aerosol with comparable accuracy to numerical parcel 465 models, which was shown to be an underlying reason for biases from ARG (Morales Betancourt 466 et al., 2014).

467 **4. Results and Discussions** 

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### 469 **4.1.Sensitivity Simulations with VBS Treatments Coupled with CB05**

471 As listed in Table 4, a number of sensitivity simulations are designed to identify the best model 472 configuration for OA treatments with the closest agreement to observations over CONUS. Those 473 sensitivity simulations consider (i) two SOA modules (MADE/SORGAM vs MADE/VBS), (ii) 474 two types of VBS treatment for POA (nonvolatile POA <u>vs.versus</u> semivolatile POA), (iii) two 475  $\Delta$ H<sub>vap</sub> treatments (default <u>vs.versus</u> the semi-empirical  $\Delta$ H<sub>vap</sub> equation by Epstein et al. (2010)), 476 (iv) three different percentages of functionalization and fragmentation (FF) (10%, 25%, and 50%), 477 (v) three sets of POA emissions (default <del>vs.</del>versus 1.5 or 3 times the original NEI POA emissions), 478 (vi) three different gas-phase mechanisms (CB05, CB6, and SAPRC07), and (vii) two different 479 aerosol activation schemes (ARG00 vs.versus combinations of different aerosol activation 480 schemes of the FN05 series: FN05, FN05/BA10, and MN14) All simulations except for CB05-481 SORG-DH contain the VBS treatments for OA. CB05-SORG-DH and CB05-VBS-DH treat POA 482 emissions as nonvolatile. In addition, the impact of two different cumulus parameterization 483 schemes: Grell – Freitas (Grell and Freitas, 2014) and the Multi-scale Kain Fritsch (MSKF) (Zheng 484 et al., 2016) scheme were also tested.

485 Table 5 summarizes the main statistics for all sensitivity simulations in terms of mean obs, 486 mean sim, Corr, NMB, and NME for hourly OC and TC concentrations against IMPROVE and 487 hourly TC concentrations against STN, respectively, over the whole CONUS domain. Figure 1 488 compares the domain mean hourly averaged observed OC or TC concentrations based on 489 IMPROVE and STN with simulated concentrations calculated based on the ratios of OA/OC 1.4 490 and 2.1 for each sensitivity simulation. The domain mean hourly averaged obs OC concentration is 0.88 µg m<sup>-3</sup> for IMPROVE, and the domain mean hourly averaged obs TC concentration is 1.03 491  $\mu$ g m<sup>-3</sup> for IMPROVE and 2.71  $\mu$ g m<sup>-3</sup> for STN. As shown in Figure 1, the simulation 492 493 CB05 SORG DH with the default SOA module SORG largely underpredicts OC and TC with the 494 largest NMBs and NMEs and the lowest Corr as compared to all other simulations with a SOA 495 module based on the VBS method. The remaining VBS simulations significantly reduce the biases 496 and errors in OC and TC from CB05\_SORG\_DH and also improve the correlation. Compared to 497 CB05 SORG DH, CB05 VBS DH with nonvolatile POA seems to perform relatively well in 498 terms of NMBs and Corr against IMPROVE OC, IMPROVE TC, and STN TC.

499 Adding the semivolatile POA treatment with 1.5 times the NEI POA emissions 500 (CB05\_POA\_DH) reduces simulated OC and TC concentrations as compared to CB05\_VBS\_DH, 501 due to the loss of mass from the semivolatile POA. As the POA mass is reduced, less surface area 502 is available for SOA precursors to condense onto, resulting in decreased OA (thus decreased OC 503 and TC) for CB05\_POA\_DH. Using the semi-empirical correlation of Epstein et al. (2010) for 504  $\Delta H_{vap}$  increases the OC and TC concentrations (CB05\_POA <u>vs.versus</u> CB05\_POA\_DH). Compared to the default  $\Delta H_{vap}$  of 30 kJ mol<sup>-1</sup> used in CB05\_POA\_DH, the semi-empirical 505 506 correlation of Epstein et al. (2010) gives much higher  $\Delta H_{vap}$  values, resulting in more of the 507 organic vapors in the particulate phase than in the gas phase. Compared to CB05-POA, the 508 simulations with various FF treatments decrease the OA concentrations, as part of the OA mass is 509 fragmented to higher volatility bins. The 10% FF case (CB05\_10% FF) does not differ significantly 510 from the no FF case (CB05\_POA). However, increasing the percentage of FF (from 10% to 25%, 511 then to 50%) decreases the OA concentrations. The FF treatments, however, even if they are more 512 representative of actual SOA atmospheric formation processes, reduce the Corr slightly (compared 513 to the cases CB05\_POA and CB05-10%FF). By doubling the POA emissions (from 1.5 to 3.0 514 times the original POA NEI emissions) for the 25% FF case (CB05\_FF25%\_EM3), the predicted 515 OC and TC concentrations are closer to the observations. When evaluated against IMPROVE OC, 516 IMPROVE TC, and STN TC, among for simulations using CB05, the simulations 517 CB05\_VBS\_DH, CB05\_POA, CB05\_FF10%, and CB05\_FF25%\_EM3 perform better than other 518 cases. The differences in the OC and TC predictions from the simulations with different gas-phase 519 mechanisms will be discussed later in Section 2.

520 Figure 2 shows the spatial distributions of simulated OC and TC concentrations overlaid with 521 observed OC from IMPROVE and TC from STN for the case CB05\_25%FF\_EM3 for the two

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522 OA/OC ratios. The model performs much better for IMPROVE OC with an OA/OC ratio of 2.1 523 as compared to 1.4, especially over eastern U.S. where the use of an OA/OC ratio of 1.4 results in 524 large overpredictions. However over the central U.S. and parts of the western U.S., the use of an 525 OA/OC ratio of 1.4 shows slightly better predictions of IMPROVE OC compared to the use of 526 OA/OC ratio of 2.1 that gives underpredictions of OC. On the other hand, the model performs 527 better for STN OC with an OA/OC ratio of 1.4 as compared to 2.1. The use of an OA/OC ratio of 528 1.4 gives better agreement with STN TC over eastern U.S. where the use of an OA/OC ratio of 2.1 529 results in large underpredictions of TC. Evaluation of OC and TC against IMPROVE and STN, 530 respectively, therefore depends heavily on the OA/OC ratio, which is site-specific. Therefore in 531 more rural sites (IMPROVE), the OA/OC ratio is more likely to be high ( $\sim 2.1$ ) with more 532 oxygenated OA, while in more urban sites (STN), the OA/OC ratio is more likely to be lower 533  $(\sim 1.4)$  due to fresher emissions and less oxidized species.

534 Figure 3 shows the POA/OA ratios for six sensitivity simulations. As mentioned earlier, the 535 observed ratio of POA/total OA is approximately 15% to 40% during the summer period over 536 various locations in the CONUS. As SOA concentrations from field campaigns are sparse at 537 different locations and at different time periods, the POA/OA ratio is used to evaluate the model's 538 capability to reproduce POA and SOA concentrations. The simulation CB05\_SORG\_DH with 539 default SORGAM SOA module largely overpredicts the POA/OA ratio, due to significant 540 underpredictions of SOA. The simulations CB05\_VBS\_DH, CB05 50%FF, and 541 CB05\_25%FF\_EM3 with various VBS treatments all have POA/OA ratios that fall within the 542 range of 0.15 to 0.4, with lower POA/OA ratios over more rural areas and higher POA/OA ratios 543 over urban areas. CB05\_VBS\_DH, however, might give too high POA concentrations over the 544 western portion of the domain as it does not consider POA to be semivolatile. Considering

545 semivolatile POA, however, without considering the fragmentation and functionalization 546 processes in the simulation CB05 POA results in too low POA/OA ratio (< 0.1 over most areas). 547 Similarly, the CB05 FF25% case also results in a large portion of CONUS with POA/OA ratios 548 of < 0.1, due to the loss of POA mass. CB05 FF50%, however, predicts reasonable POA/OA ratios, even with fragmentation/functionalization due to balanced loss of both POA and SOA mass 549 550 through fragmentation to higher volatility bins. The simulation CB05\_FF25%\_EM3 also improves 551 from CB05 FF25% by increasing the POA mass contributing to higher POA/OA ratios. 552 Figure 4 shows the observed and simulated temporal variations of SOA concentrations at the

553 two CalNex sites: Bakersfield and Pasadena in CA from May to June 2010 for the simulations 554 CB05 SORM DH. CB05\_25%FF\_EM3, CB6\_25%FF\_EM3, CB05\_VBS\_DH, and 555 SAPRC07\_25%FF\_EM3. There are large underpredictions of SOA by all runs on some days (e.g. 556 May 15 - 16, June 2 - 6, June 13 - 14) likely due to missing SOA precursor emissions. Table 6 557 shows the statistics of the simulations presented in Figure 4. The results using CB6 and SAPRC07 558 gas-phase mechanisms will be discussed in section 4.2. The observed SOA was derived based on 559 the tracer method of Kleindienst et al. (2012) which contains some uncertainties, and also likely 560 contributeds to the poor correlation for most of the cases. For example, it assumes mass fraction 561 of the tracers in secondary organic carbon is the same in the field as that in the laboratory, and the 562 tracers are assumed to be inert and are unlikely to undergo oxidation in the atmosphere, which 563 might not be the case. In addition, the SOA data from the CalNex campaign only consider 564 contributions from a small number of precursors including biogenic precursors (i.e., isoprene,  $\alpha$ -565 pinene, and  $\beta$ -caryophyllene), and the anthropogenic precursors (i.e., toluene, polycyclic aromatic 566 hydrocarbons (PAHs) and methyl butenol (MBO)).

567 As shown in Figure 4 and Table 6, the simulation CB05 SORG DH with the default 568 SORGAM SOA module significantly underpredicts observed SOA concentrations at both sites. 569 The model configuration of CB05 VBS DH has been used in a number of WRF/Chem 570 simulations published in literature (e.g., Yahya et al., 2015a; Campbell et al., 2015; Wang et al., 571 2015a, b). At Bakersfield, the simulation CB05\_VBS\_DH overpredicts the SOA concentrations 572 for almost all the days. The simulation CB05\_25%FF\_EM3, however, underpredicts the SOA 573 concentrations at Bakersfield, especially in June. The CB05 25% FF EM3 case also shows low 574 SOA concentrations throughout May and June, without much variability in SOA concentrations, 575 likely due to underestimations of original POA emissions at Bakersfield. As the S/IVOC emissions 576 for CB05\_25% FF\_EM3 are a factor of 3 of the original POA emissions from NEI, if the original 577 POA emissions from NEI are underestimated, the S/IVOC emissions will be low, resulting in low SOA concentrations due to low concentrations of condensable material. At Pasadena, both 578 CB05 VBS DH and CB05 25%FF EM3 overpredict the obs SOA from May 15<sup>th</sup> to May 30<sup>th</sup>, 579 but are unable to capture the high SOA concentrations from 2<sup>nd</sup> to 6<sup>th</sup> June. The CB05 VBS DH 580 581 case seems to perform better than the CB05\_25%FF\_EM3 case when observed SOA 582 concentrations are high. The results from this study are consistent with those from Baker et al. 583 (2015), which showed that measured PM<sub>2.5</sub> OC at Bakersfield is largely underpredicted compared 584 to Pasadena. Baker et al. (2015), however, attributed to the underpredictions of OC at Bakersfield 585 and Pasadena mainly to primary OC predicted by the baseline model, as compared to the Aerosol 586 Mass Spectrometer measurements, suggesting that OC is mostly secondary in nature in Pasadena. 587 In addition, as mentioned earlier, the simulated SOA from WRF/Chem does not consider 588 contributions from all the SOA precursors identified by their trace compounds (e.g., the biogenic 589 precursor, b-caryophyllene, and the anthropogenic precursor MBO, are not included in 590 WRF/Chem), which can help to account for the discrepancies between the simulated and observed591 SOA concentrations.

592 4.2. Sensitivity of OA predictions to different gas-phase mechanisms

593

594 Figure 1 shows that CB05\_FF25%\_EM3 produces the highest OC and TC concentrations at 595 the IMPROVE sites, followed by CB6\_FF25%\_EM3 and SAPRC07\_FF25%\_EM3, while 596 CB6\_FF25%\_EM3 produces the highest TC concentrations at the STN sites. However, the 597 differences in domain-mean simulated OC and TC between the simulations with the three different 598 gas-phase mechanisms are small, compared to the differences in simulated OC and TC due to 599 differences in VBS treatments (e.g., nonvolatile vs. versus semivolatile POA). Figure 4 also shows 600 that there are not much differences between simulated SOA concentrations with different gas-601 phase mechanisms at Bakersfield, but larger differences are found at Pasadena. . For example, 602 SAPRC07 25%FF EM3 produces much higher SOA concentrations compared to 603 CB05\_25%FF\_EM3 and CB6\_25%FF\_EM3 at Pasadena on several days (e.g., June 6-8). Figure 604 5 shows the time series of hydroxyl radical (OH) mixing ratios as well as diurnal plots of OH and 605 hydroperoxyl radical (HO<sub>2</sub>) at Pasadena from the CalNex field campaign. The time series of HO<sub>2</sub> 606 is not shown due to irregularity of the observational data. The model is able to reproduce the 607 diurnal variation of OH radicals but significantly overpredicts the daytime and peak OH mixing 608 ratios, especially for CB05 and CB6. All gas-phase mechanisms underpredict OH mixing ratios at 609 night. Among all simulations, SAPRC07 produces the closest simulated OH mixing ratios 610 compared to CB05 and CB6 gives the largest overpredictions. Similarly, the HO<sub>2</sub> mixing ratios 611 are generally overpredicted by all gas-phase mechanisms with SAPRC07 performing the best. The 612 overpredictions in OH and HO<sub>2</sub> mixing ratios do not help explain the underpredictions of SOA for several days at Pasadena where underpredictions of VOCs may be the main cause, which isconsistent with the findings of Baker et al. (2015).

615 Figure 6 shows spatial distributions of average concentrations of oxidants including ozone 616  $(O_3)$ , OH, HO<sub>2</sub>, as well as the OA species including anthropogenic SOA (ASOA), biogenic SOA 617 (BSOA), TSOA, and POA. SAPRC07-25% FF-EM3 produces the highest domain average  $O_3$ 618 mixing ratios but the lowest domain average OH+HO<sub>2</sub> mixing ratios while CB6-25%FF-EM3 619 produces the highest domain average and maximum OH+HO<sub>2</sub> mixing ratios but the lowest domain 620 average O<sub>3</sub> mixing ratios. These findings are mostly consistent from literature. For example, 621 maximum O<sub>3</sub> and OH mixing ratios over the Los Angeles area are higher for CB6 compared to 622 CB05, which are consistent with the results from Yarwood et al. (2010). SAPRC07 also generally 623 produces higher O<sub>3</sub> mixing ratios compared to CB05. However, average O<sub>3</sub> mixing ratios from 624 CB6 are expected to be higher than CB05 (rather than lower as shown in Figure 6), according to 625 the study from Nopmongcol et al. (2012) which showed higher O<sub>3</sub> mixing ratios over Europe for 626 January and July using the Comprehensive Air Quality Model with Extensions (CAMx). CB6 is a 627 relatively new gas-phase mechanism, there are very few studies that evaluated its performance 628 over a longer period of time, e.g., for the whole summer, and over CONUS. In addition, there are 629 other uncertainties in this study. For example, the emissions for CB05 are used for CB6, the 630 additional explicit VOC species in CB6 such as benzene and acetylene are not considered, which 631 can also contribute to O<sub>3</sub> formation. In addition, most locations in the U.S. in 2010 are considered 632 to by  $NO_x$ -limited with localized VOC-limited regimes from May to September (Campbell et al., 633 2015), which means that  $O_3$  formation is more likely to depend on NO<sub>x</sub> rather than VOC 634 concentrations.

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635 Table 7 shows the statistics for maximum 1-hr and 8-hr O<sub>3</sub> mixing ratios evaluated against 636 CASTNET and AQS. CASTNET sites are mainly rural sites, while AQS consists of urban, 637 suburban, and rural sites. As expected, SAPRC07 consistently produces the highest maximum 1-638 hr and maximum 8-hr  $O_3$  mixing ratios and overpredicts at AQS sites with an NMB of ~16%. 639 However, SAPRC07 performs the best at CASTNET sites, as both CB05 and CB6 significantly 640 underpredict maximum 1-hr and maximum 8-hr O<sub>3</sub> mixing ratios. At CASTNET sites, CB6 641 performs the poorest with the largest underpredictions for both maximum 1-hr and 8-hr O<sub>3</sub> mixing 642 ratios. However, CB6 predicts higher maximum 1-hr and 8-hr O<sub>3</sub> mixing ratios at AQS sites, while 643 CB05 predicts the lowest maximum 1-hr and 8-hr  $O_3$  mixing ratios at AQS sites. It is likely that 644 CB6 predicts higher O<sub>3</sub> mixing ratios at more VOC-limited sites in urban areas, while CB05 645 predicts higher O<sub>3</sub> mixing ratios at more NO<sub>x</sub>-limited areas, due to the improvement in VOC 646 speciation in CB6 compared to CB05. Overall, however, CB05 has the highest Corr and the lowest 647 NMEs for CASTNET maximum 1-hr and AQS maximum 1-hr and 8-hr O<sub>3</sub> mixing ratios. For 648 PM2.5 concentrations, CB6 produces the best performance against IMPROVE (highest Corr, 649 lowest NMB and NME) while CB05 produces the best performance against STN (highest Corr 650 and lowest NME). All 3 cases perform poorly for PM10 against AQS, with large underpredictions 651 due to the non-consideration of the coarse mode inorganic species in MADE-VBS treatments. 652 Anthropogenic SOA (ASOA) concentrations are lower for CB6 and SAPRC07 compared to

CB05. This is likely partially due to the emissions which are mapped from CB05 to CB6 and SAPRC07. The CB05 emissions are not likely to account for all anthropogenic VOC emissions in CB6 and SAPRC07, resulting in lower ASOA concentrations for CB6 and SAPRC07 compared to CB05. Biogenic SOA (BSOA) concentrations, however, are the largest for CB6, followed by SAPRC07 and CB05. BSOA concentrations are likely the highest for CB6 due to the highest  $OH+HO_2$  mixing ratios for CB6. The more extensive VOC representation and high  $O_3$  mixing ratios for SAPRC07 also likely contribute to the high BSOA concentrations for SAPRC07 compared to CB05. However, overall, the total SOA (TSOA) and POA concentrations for all three gas-phase mechanisms do not vary much, resulting in similar OA concentrations.

662 Figures 7 and 8 show the time series of simulated versuss, observed OC from IMPROVE and 663 simulated <del>vs,</del>versus observed TC from STN at several representative sites over CONUS for the 664 different gas-phase mechanisms. In general, at IMPROVE sites, CB05 gives the highest OC 665 concentrations compared to CB6 and SAPRC07 most of the time, resulting in overpredictions of 666 OC concentrations, while CB6 and SAPRC07 perform better against IMPROVE OC. The 667 overpredictions of CB05 are likely due to overpredictions in ASOA (as CB05 produces the highest ASOA concentrations compared to CB6 and SAPRC07 as shown in Figure 6). As these sites are 668 669 located in rural locations, the dominant SOA is likely to be BSOA, or downwind ASOA from more 670 urban areas. With the exception of Death Valley NP, CA, the model performs relatively well in 671 predicting IMPROVE OC concentrations. Simulations with all three gas phase mechanisms 672 overpredict OC concentrations over several days in May in Brigantine NWR, Death Valley and 673 Swangwarter, but is able to predict several of the peaks in June. All three gas-phase mechanisms, 674 however, largely underpredict OC concentrations over Death Valley from May 21<sup>st</sup> to June 30<sup>th</sup>. 675 As the Death Valley NP is a desert, the OC at Death Valley NP is most likely due to downwind 676 OC transported from upwind locations, for which the model is not able to capture due to 677 meteorological biases such as biases in wind fields. The differences between simulation results 678 from the gas-phase mechanisms are smaller for STN TC compared to IMPROVE OC, probably 679 due to similar elemental carbon (EC) concentrations for all gas-phase mechanisms, which can form 680 a significant percentage of TC. In general, all simulations with the three gas-phase mechanisms

also show similar trends (peaks and troughs) for simulated TC, likely due to influences from meteorological parameters such as wind and precipitation. Overall, all three simulations are also able to predict the magnitude and trends of STN TC concentrations relatively well. Similarly, CB05 tends to produce the highest TC concentrations, however, CB6 also does produce the highest TC concentrations for several days, for example, for some days in May in Washington, DC and Tampa, FL, as well as in June in Liberty, KS, likely due to influences of BSOA where CB6 produces the highest concentrations as shown in Figure 6.

### 688 689

### **4.3. Impact of Different VBS treatments on CDNC**

690 Table 7 shows the statistics for model evaluation for simulated CDNC against MODIS-derived 691 CDNC from Bennartz (2007) for May to June 2010. All simulations underpredict CDNC due likely 692 to underpredictions in PM and CCN concentrations and uncertainties and/or assumptions in the 693 derived CDNC based on MODIS retrievals of cloud properties by Bennartz (2007) (Zhang et al., 694 2015). For example, Bennartz (2007) derived the CDNC from cloud optical depths and cloud 695 effective radius assuming adiabatically-stratified clouds. Among all simulations with CB05, 696 CB05\_SORG\_DH produces the lowest CDNC due to underestimated OA concentrations. 697 Increasing the OA concentrations helps to reduce the negative biases for CDNC. There are small 698 differences, however, among simulated CDNC with different VBS treatments for CB05 in CDNC 699 predictions, with similar Corr ~ 0.29, NMBs of ~-29% to -27% and NMEs of ~ 47%. Figure 9 700 shows the spatial differences in predictions in warm clouds between the several simulations and 701 the simulation CB05\_VBS\_DH. CB05\_SORGAM\_DH gives the lower CDNC than 702 CB05 VBS DH, indicating that the VBS treatment in CB05 VBS DH helps to increase CDNC 703 significantly. While other simulations with semivolatile POA treatments further increase domain 704 average CDNC comparing to CB05\_VBS\_DH, the differences between CDNC predictions from

705 those simulations and CB05\_VBS\_DH are quite similar. In general, CDNC with the semivolatile 706 POA cases are higher over western U.S. but lower over eastern U.S. due to decreases in column 707 OA concentrations for the semivolatile POA cases compared to CB05 VBS DH over eastern U.S. 708 The large differences in CDNC predictions, however, are found between simulations with the 709 different gas-phase mechanisms. SAPRC07\_25%FF\_EM3 has the largest negative bias (NMB of 710 -52%) compared to all other simulations with CB05 and the simulation with CB6. Figure 10 711 compares the spatial plots for CDNC predictions for simulations with different gas-phase 712 mechanisms, as well as the surface spatial plots for total OA and inorganic  $PM_{2.5}$  concentrations. 713 The simulation with SAPRC07 shows significantly lower CDNC over northeastern U.S. 714 comparing to CDNC predictions from the other two simulations. While all three simulations show 715 similar total OA concentrations, large differences are found for their total inorganic PM<sub>2.5</sub> 716 concentrations, with SAPRC07 producing the lowest domain mean and maximum total inorganic 717 PM<sub>2.5</sub> concentrations. Compared to CB05 and CB6, the lower inorganic PM<sub>2.5</sub> concentrations 718 simulated with SAPRC07 are likely due to the low OH+HO<sub>2</sub> mixing ratios for SAPRC07 as shown 719 in Figure 6, resulting in a lower PM number concentration and lower cloud condensation nuclei 720 (CCN), thus lower CDNC.

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# 1 4.4.Sensitivity Simulations for Aerosol Activation Parameterizations

Among all OA sensitivity simulations, the simulation CB05-25%FF-EM3 gives an overall best performance in terms of OC, TC, O<sub>3</sub>, PM<sub>2.5</sub>, and CDNC evaluation, it is thus selected to test various aerosol activation parameterizations. As listed in Table 4, four sensitivity simulations are designed to test the FN05 series aerosol activation parameterizations with improved treatments comparing to the default ARG00 aerosol activation parameterization. These sensitivity simulations include the default ARG00, the FN05, the combination of FN05 and BA10, and the MN14. These 729 simulations use the MSKF scheme instead of the Morrison microphysics schemes in the previous 730 SOA runs as the MSKF scheme has a better correlation with MODIS CDNC as compared to the 731 Morrison microphysics scheme. Table 8 summarizes the model evaluation results against MODIS-732 derived CDNC from Bennartz (2007). The simulation ARG00 underpredicts CDNC with an NMB 733 of -35%. The FN05 series helps to reduce the underpredictions of CDNC significantly, because 734 they in general give higher activation fractions compared to the ARG00 parameterization under 735 most atmospheric conditions (Ghan et al., 2011). The addition of BA10 to the FN05 takes into 736 account the effects of condensation on giant CCN, which reduces the CDNC predictions and leads 737 to a negligible underprediction of CDNC (with an NMB of -0.8%) compared to a slight 738 overprediction by the FN05 with an NMB of 7.1%. MN14, which revises the original population 739 splitting method in FN05 and BA10, slightly increases the CDNC to an NMB of 4.2% comparing 740 to the FN05/BA10 simulation. The trends in the predictions of FN05, BA10, and MN14 are 741 consistent with the reported bias of ~+8%, -10% and -3%, respectively, by Morales Betancourt 742 and Nenes (2014) against the CDNC concentrations simulated from the cloud parcel model. 743 However, the Corr and NME are worse with the FN05 series and MN14. The NMEs are almost 744 doubled for the FN05 series and MN14, compared to that from the default ARG00. Figure 11 745 compares the spatial distributions of the simulated CDNC in warm clouds from ARG00, FN05 746 series, and MN14 and the MODIS-derived CDNC from Bennartz (2007). As shown in Figure 11, 747 the lower Corr and higher NMEs for the FN05 series as compared to ARG00 shown in Table 9, as 748 compared to ARG00, are due to the large overpredictions over northeastern U.S. but 749 underpredictions over other parts of the domain. The simulated CDNC from the default ARG00 750 case is similar to that from Bennartz (2007) over eastern U.S., the underpredictions are mainly 751 over western U.S. and over the ocean because of the known bias when large CCN are not present

(Morales Betancourt et al., 2014). The simulations with the FN05 series increase CDNC where CCN is high, i.e., over the northeastern U.S., resulting in overpredictions in CDNC over northeastern U.S., and does not help to improve CDNC predictions over other parts of the U.S. as well as over the ocean.

756 Figure 12 compares the simulated CCN and AOD from the CB05\_25%FF\_EM3 + MN14 case 757 with those derived from the MODIS. The model largely underpredicts CCN, especially over the 758 western part of the domain, which explains the large underprediction of CDNC also over the 759 western part of the domain. Condensation of the available water vapor occurs over CCN which are 760 concentrated over northeastern U.S., resulting in overpredictions of CDNC over northeastern U.S. 761 The lack of CCN over the ocean and the western part of the domain is related to the 762 underpredictions of AOD over the same areas. This indicates biases in number (and probably mass) 763 concentrations of column PM concentrations, especially over the ocean and western U.S. PM<sub>2.5</sub> 764 and PM<sub>10</sub> observational data are available over the surface and are both underpredicted, however, 765 there are no observational data for column concentrations of  $PM_{2.5}$  and  $PM_{10}$  for evaluation. 766 Improving the spatial distribution and magnitude of emissions for PM species and precursors for 767 the model layers at the surface and above the surface can help improve AOD and CCN predictions, 768 therefore CDNC predictions.

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# 5. Summary and Conclusions

Current regional air quality models including WRF/Chem have large uncertainties in modeling OA and aerosol-cloud feedback mechanisms such as the aerosol activation process. Comparing to the traditional OA method, the VBS treatment helps to improve OA predictions by reducing the underpredictions of OA. By including a semivolatile POA treatment, using a semi-empirical formation of Epstein et al. (2010) for  $\Delta H_{vap}$ , including 25% fragmentation and functionalization 776 as well as including additional S/IVOC emissions, the VBS treatment in WRF/Chem simulates the 777 atmospheric OA formation processes more realistically and can perform relatively well in 778 predictions of OC and TC against IMPROVE and STN. POA/OA ratios for the 779 CB05 25% FF EM3 and CB05 FF50% treatments are within the range of POA/OA ratios of ~0.15 to 0.40 from literature. Compared to the simulation with default SORGAM SOA module, 780 781 the simulations with various new VBS treatments also give better agreement with observed SOA 782 at Bakersfield and Pasadena during the CalNex field campaign from May to June 2010. However, 783 biases exist in those simulations with the VBS treatments due to several possible reasons, including 784 underestimated POA emissions, underpredicted VOC concentrations, as well as differences in the 785 SOA precursors used in the model and those contributing to the observed SOA concentrations. 786 The simulations with different gas-phase mechanisms (i.e., CB05, CB6, and SAPRC07) produce 787 in general different ASOA and BSOA concentrations. SAPRC07 produces the highest O<sub>3</sub> mixing 788 ratios, while CB6 produces the lowest OH + HO<sub>2</sub> mixing ratios. CB6 also performs the best when 789 evaluated against IMPROVE PM2.5 while CB05 performs the best when evaluated against STN 790 PM2.5 concentrations. All 3 cases perform poorly against AQS PM10 evaluation. Due to the 791 significant differences between O<sub>3</sub>, OH, and HO<sub>2</sub> mixing ratios for the three gas-phase 792 mechanisms, inorganic PM concentrations vary widely, especially between the carbon bond 793 mechanisms (CB05 and CB6) and SAPRC07, resulting in significantly different predictions of 794 CDNC. The CDNC predictions do not vary much among simulations with CB05 and different 795 VBS treatments, for example, for simulations with nonvolatile <del>vs.</del>versus semivolatile POA, and 796 with and without fragmentation and functionalization treatments. The simulation with SAPRC07 797 produces the lowest CDNC compared to those with CB05 and CB6, due to the lowest inorganic 798 PM number and mass concentrations resulted from the lowest OH and HO<sub>2</sub> mixing ratios among

all simulations. CB05 gives the best performances when evaluated against CASTNET and AQS
ozone mixing ratios, STN PM<sub>2.5</sub> concentrations and against MODIS CDNC.

801 With the default ARG00 treatment in the model, in general, all simulations with VBS 802 treatments underpredict the MODIS-derived CDNC by Bennartz (2007). By including the FN05 803 series (i.e., FN05, FN05/BA10, and MN14), the underpredictions for CDNC are greatly reduced. 804 However, the correlation coefficient and errors are worse with the FN05 series, with large 805 overpredictions over the northeastern U.S., where CCN is high. The model performs poorly for 806 AOD and CCN, likely due to inaccuracies in spatial distribution and magnitudes of PM and PM 807 precursor emissions in the model layers at the surface and above the surface. The CDNC 808 predictions can be improved by improving AOD and CCN undepredictions over western U.S. and 809 over the ocean.

### 810 Code and Data Availability

The WRF/Chem v3.7.1 code used in this paper will be available upon request. The inputs including the meteorological files, meteorological initial and boundary conditions, chemical initial and boundary conditions, model setup and configuration, and the namelist setup and instructions on how to run the simulations for a 1-day test case, as well as a sample output for a 1-day test, can be provided upon request.

# 816 Acknowledgments

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This study is funded by the National Science Foundation EaSM program (AGS-1049200) at NCSU. The emissions are taken from the 2008 NEI-derived emissions for 2010 provided by the U.S. EPA, Environment Canada, and Mexican Secretariat of the Environment and Natural Resources (Secretaría de Medio Ambiente y Recursos Naturales-SEMARNAT) and National Institute of Ecology (Instituto Nacional de Ecología-INE) as part of the Air Quality Model

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823 Evaluation International Initiative (AQMEII). The authors acknowledge use of the WRF-Chem

- 824 preprocessor tool mozbc provided by the Atmospheric Chemistry Observations and Modeling Lab
- 825 (ACOM) of NCAR and the script to generate initial and boundary conditions for WRF based on
- 826 CESM results provided by Ruby Leung, PNNL. This work also used the Stampede Extreme
- 827 Science and Engineering Discovery Environment (XSEDE) high-performance computing support
- 828 which is supported by the National Science Foundation grant number ACI-1053575. The authors
- 829 also acknowledge high-performance computing support from Yellowstone (ark:/85065/d7wd3xhc)
- 830 provided by NCAR's Computational and Information Systems Laboratory, sponsored by the
- 831 National Science Foundation and Information Systems Laboratory.

# 832 **References**

- Abdul-Razzak, H., and Ghan, S.J.: A parameterization of aerosol activation 2. Multiple aerosol
  types, J. Geophys. Res., 105, D5, 6837 6844, 2000.
- Ackermann, I.J., Hass, H., Memmesheimer, M., Ebel, A., Binkowski, F.S., and Shankar, U.: Modal
  aerosol dynamics model for Europe: Development and first applications, Atmospheric
  environment, 32, No.17, 2981-2999, 1998.

Ahmadov, R., McKeen, S.A., Robinson, A.L., Bareini, R., Middlebrook, A.M., De Gouw, J.A.,
Meagher, J., Hsie, E.-Y., Edgerton, E., Shaw, S., and Trainer, M.: A volatility basis set model
for summertime secondary organic aerosols over the eastern United States in 2006, J. Geophys.
Res., 117, doi:10.1029/2011JD016831, 2012.

Aitken, A.C., DeCarlo, P.F., Kroll, Worsnop, D.R., Huffman, J.A., Docherty, K.S., Ulbrich, I.M.,
Mohr, C., Kimmel, J.R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A., Northway, M.,
Ziemann, P.J., Canagaratna, M.R., Onasch, T.B., Alfarra, M.R., Prevot, A.S.H., Dommen, J.,
Duplissy, J., Metzger, A., Baltensperger, U., and Jimenez, J.L.: O/C and OM/OC Ratios of
primary, secondary, and ambient organic aerosols with high-resolution time-of-flight aerosol
mass spectrometry, Environ. Sci. Technol., 42(12), 4478 – 4485, doi:10.1021/es703009q,
2008.

849 Atkinson, R., Baulch, D.L., Cox, R.A., Crowley, J.N., Hampson, R.F., Hynes, R.G., Jenkin, M.E., 850 Kerr, J.A., Rossi, M.J., and Troe, J.: "Evaluated kinetic and photochemical data for atmospheric chemistry - IUPAC subcommittee on gas kinetic data evaluation for atmospheric 851 chemistry." 852 July 2005 web version available from 853 http://www.iupackinetic.ch.cam.ac.uk/index.html, 2005.

- 854 Atkinson, R.A., Baulch, D.L., Cox, R.A., Crowley, J.N., Hampson, R.F., Hynes, R.G., Jenkin, M.E., Kerr, J.A., Rossi, M.J., and Troe, J.: "Evaluated kinetic and photochemical data for 855 atmospheric chemistry - IUPAC subcommittee on gas kinetic data evaluation for atmospheric 856 857 chemistry." January 3. 2010 web version available at http://www.iupac-858 kinetic.ch.cam.ac.uk/index.html, 2010.
- Baker, K.R., Carlton, A.G., Kleindienst, T.E., Offenberg, J.H., Beaver, M.R., Gentner, D.R.,
  Goldstein, A.H., Hayes, P.L., Jimenez, J.L., Gilman, J.B., de Gouw, J.A., Woody, M.C., Pye,
  H.O.T., Kelly, J.T., Lewandowski, M., Jaoui, M., Stevens, P.S., Brune, W.H., Lin, Y-H.,
  Rubitschun, C.L., and Surratt, J.D.: Gas and aerosol carbon in California: comparison of
  measurements and model predictions in Pasadena and Bakersfield, Atmos. Chem. Phys., 15,
  5243 5258, doi:10.5194/acp-15-5243-2015, 2015.
- Barahona, D., and Nenes, A.: Parameterization of cloud droplet formation in large-scale models:
  Including effects of entrainment, J. Geophys. Res., 112, D16206, doi:10.1029/2007JF008473,
  2007.
- Barahona, D., West, R.E.L., Stier, P., Romakkaniemi, S., Kokkola, H., and Nenes, A.:
  Comprehensively accounting for the effect of giant CCN in cloud activation parameterizations,
  Atmos. Chem. Phys., 10, 2467 2473, 2010.
- Bennartz, R.: Global assessment of marine boundary layer cloud droplet number concentration
  from satellite, J. Geophys. Res. Atmos., 112, D02201, doi:10.1029/2006JD007547, 2007.
- Bergstrom, R., Denier van der Gon, H.A.C., Prevot, A.S.H., Yttri, K.E., and Simpson, D.:
  Modelling of organic aerosols over Europe (2002 2007) using a volatility basis set (VBS)
  framework: application of different assumptions regarding the formation of secondary organic
  aerosol, Atmos. Chem. Phys., 12, 8499 8527, doi:10.5194/acp-12-8499-2012, 2012.
- Boucher, O., Randall, D., Artaxo, P., Bretherton, C., Feingold, G., Forster, P., Kerminen, V.-M.,
  Kondo, Y., Liao, H., Lohmann, U., Rasch, P., Satheesh, S.K., Sherwood, S., Stevens, B., and
  Zhang, X.Y.: Clouds and Aerosols. In: Climate Change 2013: The Physical Science Basis.
  Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental
  Panel on Climate Change [Stocker, T.F., D. Qin, G.-K. Plattner, M. Tignor, S.K. Allen, J.
  Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley (eds.)]. Cambridge University Press,
  Cambridge, United Kingdom and New York, NY, USA, 2013.
- Campbell, P., Zhang, Y., Yahya, K., Wang, K., Hogrefe, C., Pouliot, G., Knote, C., Hodzic, A.,
  San Jose, R., Perez, J.L., Jimenez Guerrero, P., Baro, R., and Makar, P.: A multi-model
  assessment for the 2006 and 2010 simulations under the Air Quality Model Evaluation
  International Initiative (AQMEII) phase 2 over North America: Part I. Indicators of the
  sensitivity of O<sub>3</sub> and PM<sub>2.5</sub> formation regimes, Atmos. Environ., 115, 569 586, 2015.
- Carter, W.P.L.: Implementation of the SAPRC-99 chemical mechanism into the models-3
  framework, Report to the US EPA (<u>http://www.cert.ucr.edu/~carter/pubs/s99mod3.pdf</u>), 2000.
  Last accessed, February 19, 2016.

- Carter, W.P.L.: Development of the SAPRC07 chemical mechanism, Atmos. Environ., 44, 5324
   5335, doi:10.1016/j.atmosenv.2010.01.026, 2010.
- Chan, A.W.H., Kautzman, K.E., Chhabra, P.S., Surratt, J.D., Chan, M.N., Crounse, J.D., Kurten,
  A., Wennberg, P.O., Flagan, R.C., and Seinfeld, J.H.: Secondary organic aerosol formation
  from photooxidation of naphthalene and alkylnapthalenes: implications for oxidation of
  intermediate volatility organic compounds (IVOCs), Atmos. Chem. Phys., 9, 3049-3060, 2009.
- Cleveland, M.J., Ziemba, L.D., Griffin, R.J., Dibb, J.E., Anderson, C.H., Lefer, B., and
  Rappengluck, B.: Characterization of urban aerosol using aerosol mass spectrometry and
  proton nuclear magnetic resonance spectroscopy, Atmos. Environ., 54, 511 518,
  doi:10.1016/j.atmosenv.2012.02.074, 2012.
- Donahue, N.M., Robinson, A.L., Stanier, C.O., and Pandis, S.N.: Coupled partitioning, dilution
   and chemical aging of semivolatile organics, Environ. Sci. Tech., 40, 2635 2643, 2006.
- Donahue, M.N., Robinson, A.L., and Pandis, S.N.: Atmospheric organic particulate matter: From
   smoke to secondary organic aerosol, Atmos. Environ. 43, 94-106, 2009.
- Emmons, L.K., Walters, S., Hess, P.G., Lamarque, J.-F., Pfister, G.G., Fillmore, D., Granier, C.,
  Guenter, A., Kinnison, D., Laepple, T., Orlando, J., Tie, X., Tyndall, G., Wiedinmyer, C.,
  Baughcum, S.L., and Kloster, S.: Description and evaluation of the Model for Ozone and
  Related chemical Tracers, version 4 (MOZART-4), Geosci. Model Dev., 3, 43 67, 2010.
- ENVIRON: User's guide to the comprehensive air quality model with extensions, version 6, 2013.
   Available at www.camx.com/files/camxusersguide\_v6-10.pdf
- Epstein, S.A., Riipinen, I., and Donahue, N.M.: A semiempirical correlation between enthalpy of
   vaporization and saturation concentration for organic aerosol, Environ. Sci. Technol., 44, 743
   -748, 2010.
- Farina, S.C., Adams, P.J., and Pandis, S.N.: Modeling global secondary organic aerosol formation
  and processing with the volatility basis set: Implications for anthropogenic secondary organic
  aerosol, J. Geophys. Res., 115, D09202, doi:10.1029/2009JD013046, 2010.
- Fountoukis, C., and Nenes, A.: Continued development of a cloud droplet formation
  parameterization for global climate models, J. Geophys. Res., 110, D11212,
  doi:10.1029/2004JD00591, 2005.
- Gantt, B., He, J., Zhang, X., Zhang, Y., and Nenes, A.: Incorporation of advanced aerosol activation treatments into CESM/CAM5: model evaluation and impacts on aerosol indirect effects, Atmos. Chem. Phys., 14, 7485 7497, doi:10.5194/acp-14-7485-2014, 2014.
- Ghan, S.J., Abdul-Razzak, H., Nenes, A., Ming, Y., Liu, X., Ovchinnikov, M., Shipway, B.,
  Meskhidze, N., Xu, J., and Shi, X.: Droplet nucleation: Physically-based parameterizations and
  comparative evaluation, J. Adv. Model. Earth Syst., 3, M10001, doi:10.1029/2011ms000074,
  2011.

- Glotfelty, T., He, J. and Zhang, Y.: Updated organic aerosol treatments in CESM/CAM5:
   development and initial application, in preparation, 2016.
- Gong, S., Barrie, L.A. and Blanchet, J.P.: Modeling sea salt aerosols in the atmosphere: 1. Model
  development, J. Geophys. Res., 102, 3805-3818, doi:10.1029/96JD02953, 1997.
- Grell, G.A. and Freitas, S.R.: A scale and aerosol aware stochastic convective parameterization
  for weather and air quality modeling, Atmos. Chem. Phys., 14, 5233-5250, doi:10.5914/acp14-5233-2014, 2014.
- Grieshop, A.P., Logue, P., Donahue, J.M., Robinson, A.L., Laboratory investigation of
  photochemical oxidation of organic aerosol from wood fires 1: measurement and simulation
  of organic aerosol evolution, Atmos. Chem. Phys., 9, 1263-1277, 2009.
- Guenther, A., Kart, T., Harley, P., Wiedinmyer, C., Palmer, P.I. and Geron, C.: Estimates of global
  terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and Aerosols from
  Nature), Atmos. Chem. Phys., 6, 3181-3210, 2006.
- Hallquist, M., Wenger, J.C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J.,
  Donahue, N.M., George, C., Goldstein, A.H., Hamilton, J.F., Herrman, H., Hoffman, T.,
  Iinuma, Y., Jang, M., Jenkin, M.E., Jimenez, J.L., Kiendler-Scharr, A., Maenhaut, W.,
  McFiggans, G., Mentel, T.F., Monod, A., Prevot, A.S.H., Seinfeld, J.H., Surratt, J.D.,
  Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic
  aerosol: current and emerging issues, Atmos. Chem. Phys., 9, 5155 5236, 2009.
- 947 Hayes, P.L., Ortega, A.M., Cubison, M.J., Froyd, K.D., Zhao, Y., Cliff, S.S., Hu, W.W., Toohey, 948 D.W., Flynn, J.H., Lefer, B.L., Grossberg, N., Alvarez, S., Rappengluck, B., Taylor, J.W., 949 Allan, J.D., Holloway, J.S., Gilman, J.B., Kuster, W.C., de Gouw, J.A., Massoli, P., Zhang, 950 X., Liu, J., Weber, R.J., Corrigan, A.L., Russell, L.M., Isaacman, G., Worton, D.R., Kreisberg, 951 N.M., Goldstein, A.H., Thalman, R., Waxman, E.M., Volkamer, R., Lin, Y.H., Surratt, J.D., 952 Kleindienst, T.E., Offenberg, J.H., Dusanter, S., Griffith, S., Stevens, P.S., Brioude, J., 953 Angevine, W.M., and Jimenez, J.L.: Organic aerosol composition and sources in Pasadena, 954 California during the 2010 CalNex campaign, J. Geophys. Res. Atmos., 118, 9233-9257, 955 doi:10.1002/jgrd.50530, 2013.
- He, J., and Zhang, Y.: Improvement and further development in CESM/CAM5: gas-phase
  chemistry and inorganic aerosol treatments, Atmos. Chem. Phys., 14, 9171 9200,
  doi:10.5194/acp-14-9171-2014, 2014.
- Henderson, B.H., Akhtar, F., Pye, H.O.T., Napelenok, S.L., and Hutzell, W.T.: A database and tool for boundary conditions for regional air quality modeling: description and evaluation, Geosci. Model Dev., 7, 339 360, doi:10.5194/gmd-7-339-2014, 2014.
- Hodzic, A., Jimenez, J.L., Madronich, S., Canagaratna, M.R., DeCarlo, P.F., Kleinman, L., and
   Fast, J.: Modeling organic aerosols in a megacity: potential contribution of semi-volatile and

- 964 intermediate volatility primary organic compounds to secondary organic aerosol formation,
  965 Atmos. Chem. Phys., 10, 5491-5514, 2010.
- Hong, S.-Y., Noh, Y. and Dudhia, J.: A new vertical diffusion package with an explicit treatment
   of entrainment processes, Mon. Wea. Rev., 134, 2318-2341, 2006.
- Hong, S.-Y.: A new stable boundary-layer mixing scheme and its impact on the simulated East
  Asian summer monsoon, Q.J.R. Meteorol. Soc., 136, 1481 1496, doi:0.1002/qj.665, 2010.
- 970 IUPAC: "Evaluated Kinetic and Photochemical Data". IUPAC Subcommittee on Gas Kinetic Data
   971 Evaluation for Atmospheric Chemistry. Web Version. Available at
   972 http://www.iupackinetic.ch.cam.ac.uk. Latest data sheets dated June, 2006.
- Jathar, S.H., Farina, S.C., Robinson, A.L., and Adams, P.J.: The influence of semi-volatile and
  reactive primary emissions on the abundance and properties of global organic aerosol, Atmos.
  Chem. Phys., 11, 7727 7746, doi:10.5194/acp-11-7727-2011, 2011.
- 976 Jimenez, J.L., Canagaratna, M.R., Donahue, N.M., Prevot, A.S.H., Zhang, Q., Kroll, J.H., 977 DeCarlo, P.F., Allan, J.D., Coe, H., Ng, N.L., Aiken, A.C., Docherty, K.S., Ulbrich, I.M., 978 Grieshop, A.P., Robinson, A.L., Duplissy, J., Smith, J.D., Wilson, K.R., Lanz, V.A., Hueglin, 979 C., Sun, Y.L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, 980 M., Kulmala, M., Tomlinson, J.M., Collins, D.R., Cubison, M.J., Dunlea, E.J., Huffman, J.A., 981 Onasch, T.B., Alfarra, M.R., Williams, P.I., Bower, K., Kondo, Y., Schneider, J., Drewnick, 982 F., Borrman, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., 983 Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J.Y., Zhang, Y.M., Dzepina, K., Kimmel, 984 J.R., Sueper, D., Jayne, J.T., Herndon, S.C., Trimborn, A.M., Williams, L.R., Wood, E.C., 985 Middlebrook, A.M., Kolb, C.E., Baltensperger, U., and Worsnop, D.R.: Evolution of Organic 986 Aerosols in the Atmosphere, Science, 326 (5959), 1525 – 1529, doi:10.1126/science.1180353, 987 2009.
- Jones, S. and Creighton, G.: AFWA dust emission scheme for WRF/Chem-GOCART, 2011 WRF
   workshop, June 20-24, Boulder, CO, USA, 2011.
- Kim, Y., Sartelet, K., and Seigneur, C.: Formation of secondary aerosols over Europe: comparison
  of two gas-phase mechanisms, Atmos. Chem. Phys., 11, 583 598, doi:10.5194/acp-11-5832011, 2011.
- Kleindienst, T.E., Jaoui, M., Lewandowski, M., Offenberg, J.H., and Docherty, K.S.: The formation of SOA and chemical tracer compounds from the photooxidation of naphthalene and its methyl analogs in the presence and absence of nitrogen oxides, Atmos. Chem. Phys., 8711 8726, doi:10.5194/acp-12-8711-2012, 2012.
- Laaksonen, A., Malila, J., Nenes, A., Hung, H.Mop., Chen, J.P.: Surface fractal dimension, water
  adsorption efficiency, and cloud nucleation activity of insoluble aerosol, *Nat.Sci.Rep.*, 6,
  25504, doi:10.1038/srep25504, 2016.

- Lane, T.E., Donahue, N.M., and Pandis, S.N.: Simulating secondary organic aerosol formation
  using the volatility basis-set approach in a chemical transport model, 42, 7439 7451, 2008.
- Lewandowski, M., Piletic, I.R., Kleindienst, T.E., Offenberg, J.H., Beaver, M.R., Jaoui, M.,
  Docherty, K.S., and Edney, E.O.: Secondary organic aerosol characterisation at field sites
  across the United States during the spring-summer period, Intern. J. Environ. Anal. Chem.,
  doi:10.1080/030674, 2013.
- Li, J., Zhang, H., and Ying, Q.: Comparison of the SAPRC07 and SAPRC99 photochemical mechanisms during a high ozone episode in Texas: Differences in concentrations, OH budget and relative response factors, Atmos. Environ., 25 – 35, doi:10.1016/j.atmosenv.2012.02.034, 2012.
- Luecken, D.J., Phillips, S., Sarwar, G., and Jang, C.: Effects of using the CB05 <u>vs.versus</u>
   SAPRC99 <u>vs.versus</u> CB4 chemical mechanism on model predictions: Ozone and gas-phase
   photochemical precursor concentrations, Atmos. Environ., 42, 5805 5820, 2008.
- May, A.A., Presto, A.A., Hennigan, C.J., Nguyen, N.T., Gordon, T.D., and Robinson, A.L.: Gasparticle partitioning of primary organic aerosol emissions: (1) Gasoline vehicle exhaust,
  Atmos. Environ., 77, 128 139, 2013a.
- May, A.A., Presto, A.A., Hennigan, C.J., Nguyen, N.T., Gordon, T.D., and Robinson, A.L.: Gasparticle partitioning of primary organic aerosol emissions: (2) Diesel vehicles, Environ. Sci.
  Tech., 47, 8288 8296, 2013b.
- May, A.A., Levin, E.J.T., Hennigan, C.J., Riipinen, I., Lee, T., Collett Jr., J.R., Jimenez, J.L.,
  Kreidenweis, S.M., and Robinson, A.L.: Gas-particle partitioning of primary organic aerosol
  emissions, 3. Biomass burning, J. Geophys. Res., 118, 11327 11338, doi:10.1002/jgrd.50828,
  2013c.
- Morales Betancourt, R., and Nenes, A.: Droplet activation parameterization: the populationsplitting concept revisited, Geosci. Model Dev., 7, 2345 2357, doi:10.5194/gmd-7-23452014, 2014.
- Morrison, H., Thompson, G. and Tatarskii, V.: Impact of cloud microphysics on the development
   of trailing stratiform precipitation in a simulated squall line: Comparison of One- and Two Moment Schemes, Mon. Wea. Rev., 137, 991-1007, 2009.
- Murphy, B.N., and Pandis, S.N.: Simulating the formation of semivolatile primary and secondary
   organic aerosol in a regional chemical transport model, Environ. Sci. Technol., 2009, 43, 4722
   4728, 2009.
- Myhre, G., Shindell, D., Breon, F.-M., Collins, W., Fuglestvedt, F., Huang, J., Koch, D.,
  Lamarque, J.-F., Lee, D., Mendoza, B., Nakajima, T., Robock, A., Stephens, G., Takemura,
  T., and Zhan, H.: Anthropogenic and Natural Radiative Forcing in: Climate Change 2013: The
  Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of

- the Intergovernmental Panel on Climate Change [Stocker, T.F., D. Qin, G.-K. Plattner, M.
  Tignor, S.K. Allen, J. Boschung, A. Nauels, Y. Xia, V. Bex and P.M. Midgley (eds.)].
  Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2013.
- Nenes, A. Ghan, S., Abdul-Razzak, H., Chuang, P.Y., and Seinfeld, J.: Kinetic limitations on cloud
   droplet formation and impact on cloud albedo, Tellus (2001), 53B, 133 149, 2001.
- Nopmongcol, U., Koo, B., Tai, E., Jung, J., Piyachaturawat, P., Emery, C., Yarwood, G., Pirovano,
  G., Mitsakou, C., and Kallos, G.: Modeling Europe with CAMx for the Air Quality Model
  Evaluation International Initiative (AQMEII), Atmos. Environ., 53, 177 185, 2001.
- Odum, J.R., Hoffman, T., Bowman, F., Collins, D., Flagan, R.C., and Seinfeld, J.H.: Gas/Particle
  Partitioning and Secondary Organic Aerosol Yields, Environ. Sci. Tech., 30(8), 2580 2585,
  doi:10.1021/es950943+, 1996.
- Pouliot, G., van der Gon, H.D., Kuenen, J., Makar, P., Zhang, J., and Moran, M.: Analysis of the
  emission inventories and model-ready emission datasets for Europe and North America for
  phase 2 of the AQMEII project, Atmos. Environ., 115, 345 360, 2015.
- Pye, H. and Seinfeld, J.H.: A global perspective on aerosol from low-volatility organic compounds,
  Atmos. Chem. Phys., 10, 4377 4401, doi:10.5194/acp-10-4377-2010, 2010.
- 1052Rao, V., Tooly, L., and Drukenbrod, J.: 2008 National Emissions Inventory: Review, Analysis and1053Highlights, EPA-454/R-13-005, accessed online at1054http://www.epa.gov/ttn/chief/net/2008report.pdf, 2013, last access October 9th, 2015.
- Sander, S.P., Friedl, R.R., Golden, D.M., Kurylo, M.J., Huie, R.E., Orkin, V.L., Moortgat, G.K.,
  Ravishankara, A.R., Kolb, C.E., Molina, M.J., and Finlayson-Pitts, B.J.: Chemical Kinetics
  and Photochemical Data for use in Atmospheric Studies, Evaluation Number 14. NASA Jet
  Propulsion Laboratory. February. Available from
  http://jpldataeval.jpl.nasa.gov/download.html, 2013.
- Sander, S.P., Friedl, R.R., Golden, D.M., Kurylo, M.J., Moortgat, G.K., Wine, P.H., 1060 1061 Ravishankara, A.R., Kolb, C.E., Molina, M.J., Finlayson-Pitts, B.J., Huie, R.E., and Orkin, 1062 V.L.: "Chemical Kinetics and Photochemical Data for use in Atmospheric Studies, Evaluation 1063 Number 15. NASA Jet Propulsion Laboratory." July. Available from 1064 http://jpldataeval.jpl.nasa.gov/download.html, 2006.
- Sarwar, G., Fahey, K., Napelenok, S., Roselle, S. and Mathur, R.: Examining the impact of CMAQ
   model updates on aerosol sulfate predictions, the 10<sup>th</sup> Annual CMAS Models-3 User's
   Conference, October, Chapel Hill, NC, 2011.
- Schell B., Ackermann, I.J., Hass, H., Binkowski, F.S., and Ebel, A.: Modeling the formation of
  secondary organic aerosol within a comprehensive air quality model system, J. Geophys.
  Res. 106, 28275-28293, 2001.

- Shearer, S.M., Harley, R.A., Jin, L., and Brown, N.J.: Comparison of SAPRC99 and SAPRC07
  mechanisms in photochemical modeling for central California, Atmos. Environ., 46, 205 –
  216, doi:10.1016/j.atmosenv.2011.09.079, 2012.
- Shrivastava, M.K., Lane, T.E., Donahue, N.M., Pandis, S.N., and Robinson, A.L.: Effects of gas
   particle partitioning and aging of primary emissions on urban and regional organic aerosol
   concentrations, J. Geophys. Res., 113, D18301, doi:10.1029/2007JD009735, 2008.
- Shrivastava, M., Fast, J., Easter, R., Gustafson Jr., W.I., Zaveri, R.A., Jimenez, J.L., Saide, P. and
  Hodzic, A.: Modeling organic aerosols in a megacity: comparison of simple and complex
  representations of the volatility basis set approach, Atmos. Chem. Phys., 11, 6639 6662,
  doi:10.5194/acp-11-6639-2011, 2011.
- Shrivastava, M., Zelenyuk, A., Imre, D., Easter, R., Beranek, J., Zaveri, R.A., and Fast, J.:
  Implications of low volatility SOA and gas-phase fragmentation reactions on SOA loadings and their spatial and temporal evolution in the atmosphere, J. Geophys. Res. Atmos., 118, 3328 - 3342, doi:10.1002/jgrd.50160, 2013.
- Shrivastava, M., Easter, R.C., Liu, X., Zelenyuk, A., Singh, B., Zhang, K., Ma, P.-L., Chand, D.,
  Ghan, S., Jimenez, J.I., Zhang, Q., Fast, J., Rasch, P.J., and Titta, P.: Global transformation
  and fate of SOA: Implications of low-volatility SOA and gas-phase fragmentation reactions,
  J. Geophys. Res. Atmos., 120, 4169 4195, doi:10.1002/2014JD022563, 2015.
- 1089 Stockwell, W., Kirchner, F., Kuhn, M., and Seefeld, S.: A new mechanism for regional 1090 atmospheric chemistry modeling, J. Geophys. Res., 102, D22, 25847 – 25879, 1997.
- Tsimpidi, A.P., Karydis, V.A., Zavala, M., Lei, W., Molina, L., Ulbrich, I.M., Jimenez, J.L., and
  Pandis, S.N.: Evaluation of the volatility basis-set approach for the simulation of organic
  aerosol formation in the Mexico City metropolitan area. Atmospheric Chemistry and Physics
  10, 525-546, 2010.
- Turpin, B.J., and Lim, H.-J.: Species Contributions to PM2.5 Mass concentrations: Revisiting
   Common Assumptions for Estimating Organic Mass, Aero. Sci. Technol., 35, 1, 602 610,
   doi:10.1080/02786820119445, 2001.
- 1098US EPA: 2008 National Emissions Inventory, version 3 Technical Support Document, September10992013, Draft, accessed online at1100<a href="http://www3.epa.gov/ttn/chief/net/2008neiv3/2008\_neiv3\_tsd\_draft.pdf">http://www3.epa.gov/ttn/chief/net/2008neiv3/2008\_neiv3\_tsd\_draft.pdf</a>, 2013, last accessed1101October 10<sup>th</sup>, 2015.
- Wang, K., and Zhang, Y.: Application, evaluation, and process analysis of the US EPA's 2002
  Multiple-Pollutant Air Quality Modeling Platform, Atmos. And Clim. Sci., 2, 254 289, 2012.
- Wang, K., Zhang, Y., Yahya, K., Wu, S.-Y., and Grell, G.: Implementation and initial application
   of new chemistry-aerosol options in WRF/Chem for simulating secondary organic aerosols

 1106
 and aerosol indirect effects, Atmos. Environ., 115, 716 - 723,

 1107
 doi:10.1016/j.atmosenv.2015.12.007, 2015.

1108 Xu, L., Guo, H., Boyd, C.M., Klein, M., Bougiatioti, A., Cerully, K.M., Hite, J.R., Isaacman-VanWertz, G., Kreisberg, N.M., Knote, C., Olson, K., Koss, A., Goldstein, A.H., Hering, S.V., 1109 1110 de Gouw, J., Baumann, K., Lee, S.-H., Nenes, A., Weber, R.J., and Ng, N.L.: Effects of anthropogenic emissions on aerosol formation from isoprene and monoterpenes in the 1111 1112 southeastern United States. Proc. Natl. Acad. Sci.. 112. 1. 37 42. \_ 1113 doi:10.1073/pnas.1417609112, 2015.

- Yahya, K., Wang, K., Campbell, P., Glotfelty, T., He, J., and Zhang, Y.: Decadal evaluation of
  regional climate, air quality and their interactions over the continental US using WRF/Chem
  version 3.6.1, Geosci. Model Dev., 9, 671 695, doi:10.5194/gmd-9-671-2016, 2016a.
- Yahya, K., Campbell, P., Chen, Y., Glotfelty, T., Pirhalla, M., and Zhang, Y.: Downscaling CESM
  using WRF/Chem: Decadal Application for Regional Air Quality and Climate Modeling over
  the U.S. under the Representative Concentration Pathways Scenarios, in preparation, 2016b.
- Yarwood, G., Rao, S., Yocke, M., and Whitten, G.Z.: Final Report Updates to the Carbon Bond
  Chemical Mechanism: CB05, Rep. RT-04-00675, 246 pp., Yocke and Co., Novato, CA, 2005.
- Yarwood, G., Whitten, G.Z., Jung, J., Heo, G., and Allen, D.T.: Final Report Development,
  Evaluation and Testing of Version 6 of the Carbon Bond Chemical Mechanism (CB6): Work
  Order No. 582-7-84005-FY10-26, ENVIRON, Novato, CA, 2010.
- Yu, S., Eder, B., Dennis, R., Chu, S.-H., and Schwartz, S.: New unbiased symmetric metrics for
  evaluation of air quality models, Atmos. Sci. Lett., 7, 26 34, 2006.
- Zhang, Q., Worsnop, D.R., Canagaratna, M.R., and Jimenez, J.L.: Hydrocarbon-like and oxygenated organic aerosols in Pittsburgh: insight into sources and processes of organic aerosols, Atmos. Chem. Phys., 5, 3289 3311, doi:10.5194/acp-5-3289-2005, 2005.
- 1130 Zhang, Q., Jimenez, J.L., Canagaratna, M.R., Allan, J.D., Coe, H., Ulbrich, I., Alfarra, M.R., Takami, A., Middlebrook, A.M., Sun, Y.L., Dzepina, K., Dunlea, E., Docherty, K., DeCarlo, 1131 1132 P.F., Salcedo, D., Onasch, T., Jayne, J.T., Miyoshi, T., Shimono, A., Hatakeyama, S., 1133 Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, 1134 K., Williams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R.J., Rautiainen, J., Sun, J.Y., Zhang, Y.M., and Worsnop, D.R.: Ubiquity and dominance of oxygenated species in organic 1135 aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, Geophys. Res. 1136 Lett., 34, L13801, doi:10.1029/2007GL029979, 2007. 1137
- Zhang, Q., Jimenez, J.L., Canagaratna, M.R., Ulbrich, I.M., Ng, N.L., Worsnop, D.R., and Sun,
  Y.: Understanding atmospheric organic aerosols via factor analysis of aerosol mass
  spectrometry: a review, Anal. Bioanal. Chem., 401(10), 3045 3067, doi:10.1007/s00216011-5355-y, 2011.

- Zhang, Y., Chen, Y.-C., Sarwar, G., and Schere, K.: Impact of gas-phase mechanisms on
  WRF/Chem predictions: Mechanism implementation and comparative evaluation, J. Geophys.
  Res., 117, D01301, doi:10.1029/2011JD015775, 2012.
- Zhang, Y., Zhang, X., Wang, K., He, J., Leung, L.R., Fan, J., and Nenes, A.: Incorporating an advanced aerosol activation parameterization into WRF-CAM5: Model evaluation and parameterization intercomparison, J. Geophys. Res. Atmos., 120, doi:10.1002/2014JD023051, 2015.
- Zhao, Y., Hennigan, C.J., May, A.A., Tkacik, D.S., de Gouw, J.A., Gilman, J.B., Kuster, W.C.,
  Borbon, A., and Robinson, A.L.: Intermediate-volatility organic compounds: A large source
  of secondary organic aerosol, Environ. Sci. Technol., 48, 13743 13750,
  doi:10.1021/es5035188, 2014.
- Zheng, Y., Alapaty, K.A., Herwehe, J.A., Del Genio, A.D., and Niyogi, D.: Improving highresolution weather forecasts using the Weather Research and Forecasting (WRF) Model with
  an updated Kain-Fritsch scheme, Mon. Wea. Rev., 144, 833 860, doi:10.1175/mwr-d-150005.1, 2016.
- 1157

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# Tables

Source	Ahmadov et al., 2012	Shrivastava et al., 2011	Farina et al., 2010	Jathar et al., 2011	Hodzic et al., 2010
Model	WRF-Chem	WRF-Chem	GISS II' GCM	GISS II' GCM	CHIMERE
Domain	CONUS	Mexico City	Global	Global	Mexico City
Spatial Resolution	20 km and 60 km	Nested 3 km within 12 km	$4^{\circ} \times 5^{\circ}$	$4^{\circ} \times 5^{\circ}$	$5 \text{ km} \times 5 \text{ km}$ and $35 \text{ km} \times 35 \text{ km}$
Emissions of SVOCs, IVOCs, and VOCs	Only VOCs; no emissions of SVOCs and IVOCs	SVOC emissions 3 times POA emissions for both anthropogenic and biomass burning emissions. IVOC emissions 1.5 times POA emissions	POA is treated as nonvolatile and nonreactive, but acts as absorbing phase for SOA condensation, forming 1 OA phase	SVOC emissions are represented by the traditional emission inventory. IVOC emissions are 1.5 times traditional emissions	SVOC emissions 3 times POA emissions. IVOC emissions 1.5 times POA emissions
No. of VBS bins	4	2 and 9	4	9	9
Aging	Yes and No Simulations with aging: each oxidation step produces 7.5% additional mass	Yes and No Simulations with aging: each oxidation step produces 15% additional mass	Yes	Yes and No. Each oxidation step does not produce any additional mass	Yes. 2 cases below: (i)Each oxidation step produces 7.5% additional mass (ii)Each oxidation step reduces the volatility by 2 orders of magnitude and 40% of additional mass produced
Observational data	SEARCH, STN, IMPROVE	MILAGRO 2006 field campaign	IMPROVE, EMEP	IMPROVE, FAME, MILAGRO, SOAR	MILAGRO 2006 field campaign
Variables evaluated	OC and EC	OA, HOA, OOA O:C ratio	OM:OC of 1.8	OA (surface), HOA, OOA	HOA, OOA, BBOA, O:C ratio
Summary of results with VBS framework with/without aging compared to the traditional SOA approach	-Improved diurnal variability -Results without the aging process underestimate OC throughout the day	-HOA and OOA: Lower negative bias with addition of S/IVOC emissions -OOA: 2 bin VBS better results than 9 bin VBS -Underprediction of O:C ratio in both cases	-IMPROVE: improved with aging -EMEP: aging further biases already high OA predictions	-Adding IVOC emissions improves performance, however underprediction remains in winter months	-HOA overpredicted during nighttime -Case(i):Modeled O:C 3 times lower than observed -Case(ii):Better agreement for O:C but SOA generally overestimated

Table 1. Summery of covered literature studies of VRS treatments in verious regional and alc	
-1 and $-1$ . Notice that $v$ is several inclaimed Number of $v$ is a negative of values revealed and $v$	
Tudio 1. Dummary of beyend monature studies of y DD freatments in yarrous regional and gr	Jour mouchs.

Note: HOA: Hydrocarbon-like OA – Reduced specie of OA, generally understood as a surrogate for urban combustion-related POA; OOA: Oxygenated OA – Characterized by its high oxygen content and generally understood as a surrogate for SOA; BBOA: Biomass burning OA

Source	Bergstrom et al., 2012	Lane et al., 2008	Donahue et al., 2009	Murphy et al., 2009	This work
Model	EMEP	PMCAMx	PMCAMx	PMCAMx	WRF/Chem
Domain	Europe, a large part of	Eastern U.S.	Eastern U.S.	Eastern U.S.	CONUS with parts of
	the North Atlantic and				Canada and Mexico
	Arctic areas				
Spatial Resolution	$50 \text{ km} \times 50 \text{ km}$	$36 \text{ km} \times 36 \text{ km}$	NA	$36 \text{ km} \times 36 \text{ km}$	$36 \text{ km} \times 36 \text{ km}$
Emissions of SVOCs,	VOCs are present.	Only VOCs; SVOCs	Additional IVOCs	IVOC emissions are 0.2	S/IVOCs are 1.5 to 3
IVOCs and VOCs	S/IVOCs are 2.5 times	and IVOCs not added	added but details are	to 0.8 times the	times the nonvolatile
	the POA emissions		not given	nonvolatile POA	POA NEI emissions
				emission rates	
No. of VBS bins	4 for SOA components	4	9	10	4 for SOA components
	and 9 for POA				and 9 for POA
Aging	Yes and No.	Yes.	Yes.	Yes.	Yes.
	Each oxidation step	No additional mass	No additional mass	No additional mass	Each oxidation step
	produces 7.5%	produced for each	produced for each	produced for each	produces 7.5%
	additional mass	oxidation step.	oxidation step.	oxidation step.	additional mass
Observational data	CARBOSOL, SORGA,	STN, IMPROVE	NA	STN, IMPROVE	STN, IMPROVE, field
	Gote-2005				data
Variables evaluated	TC, OC	OA	POA, OPOA, SOA	OA	TC, OC, POA/OA
Summary of results	-Addition of aging	-Addition of aging	-Aging results in better	-Slight overprediction	-Large improvements in
with VBS framework	reactions improve	reactions overpredicts	model predictions	with IMPROVE	predictions
with/without aging	summertime results but	the OA concentrations		-Underprediction with	
compared to the	has little or negative	in rural IMPROVE		STN	
traditional SOA	consequences in	stations but improves			
approach	wintertime	the model performance			
	-Deteriorations of	in urban areas			
	model results with				
	increased aging at				
	urban influenced sites				
	in southern Europe				

Table 1. (cont). Summary of several literature studies of VBS treatments in various regional and global models.

Note: TC: total carbon; OC: Organic carbon; OPOA: oxidized POA

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	CB05-Cl <sup>1</sup>	CB6	SAPRC07 <sup>2</sup>
No. of species	70	114	118
No. of reactions	156	218	599
Lumping method	Lumped structure based on carbon bonds	Lumped structure based on carbon bonds	Lumped species based on their reactivity towards hydroxyl (OH)
Kinetic Data for rate constants	Mostly from IUPAC (Atkinson et al., 2005). NASA/JPL (Sander et al., 2003) values were used in some cases where IUPAC data was not available.	New information from IUPAC (Atkinson et al., 2010) and NASA (Sander et al., 2006)	Mainly from IUPAC (2006) and NASA (Sander et al., 2006).
Photolysis data	Mainly from SAPRC99 chemical mechanism. IUPAC (Atkinson et al., 2005) was used if it differs significantly from SAPRC99.	New information from IUPAC (Atkinson et al., 2010) and NASA (Sander et al., 2006)	Mainly from IUPAC (2006) and NASA (Sander et al., 2006).
Ozone chemistry	Slightly underpredict O <sub>3</sub> mixing ratios with isoprene and in synthetic urban mixtures in chamber experiments.	Reduced underprediction in $O_3$ mixing ratios from benzene, toluene, and xylene, but forms $O_3$ from isoprene too slowly compared to CB05.	Slightly underpredict O <sub>3</sub> mixing ratios at low NO <sub>x</sub> levels in chamber experiments.
Organic nitrate	2 reactions involving organic nitrate (NTR).	Additional NO <sub>x</sub> recycling from organic nitrate to represent fate of NO <sub>x</sub> over multiple days.	Added peroxy+NO reactions to form organic nitrate.
Chlorine chemistry	20 additional reactions for Cl chemistry involving species Cl <sub>2</sub> , HOCl, Cl, ClO, and FMCl.	CB05 chlorine chemistry included in this work.	22 base chlorine reactions involving CL <sub>2</sub> , CLNO, CLONO, CLNO <sub>2</sub> , CLONO <sub>2</sub> , HOCl, and 26 additional reactions involving organic products
Organic chemistry	- Explicit organic aerosol precursors, e.g., isoprene, toluene, xylene, α-pinene, β- pinene.	<ul> <li>Explicit long-lived and abundant organic compounds including propane, acetone, benzene and acetylene added</li> <li>Extensive revision of isoprene and aromatics chemistry</li> <li>Formation of alpha-dicarbonyl compounds (glyoxal, glycoaldehyde, methylglyoxal)</li> <li>Updates to peroxy radical chemistry that will improve formation of H<sub>2</sub>O<sub>2</sub></li> </ul>	<ul> <li>Reformulated reactions of peroxy radicals so that effects of changes in NO<sub>x</sub> conditions on organic product formation is more accurately represented</li> <li>Most comprehensive representation of VOCs compared to other gas-phase mechanisms</li> </ul>
3-D host models	Implemented into WRF/Chem v3.6.1 by Wang et al. (2014). Also available in WRF/Chem v3.7.1	Implemented in CAMx by ENVIRON (2013)	Implemented in CMAQ (Carter, 2010)
Reference	Yarwood et al. (2005)	Yarwood et al. (2010) ENVIRON (2013)	Carter (2010)

Table 2. Summary of main characteristics of CB05, CB6, and SAPRC07 gas-phase mechanisms.

<sup>1</sup>CB05 gas-phase mechanism with reactive chlorine chemistry (Yarwood et al., 2005) <sup>2</sup>SAPRC07 uncondensed and expanded version C, which includes reactions for peroxy radical operators (Carter, 2010).

		,	,	
Log Ci* at	Normalized	Fraction for	Fraction for	New calculated
298K	fraction for	gasoline	biomass	fraction for all
	stationary	emissions from	burning	sources based
	emissions based	May et al.	emissions from	on Shrivastava
	on	(2013a)	May et al.	et al. (2011),
	anthropogenic		(2013c)	May et al.
	emissions from			(2013a, c), and
	Shrivastava et			% distribution
	al. (2011)			of NEI
				emissions
-2	0.04	0.14	0.2	0.1754
-1	0.02	0.13	0.0	0.0141
0	0.03	0.15	0.1	0.0961
1	0.05	0.26	0.1	0.1084
2	0.07	0.15	0.2	0.1799
3	0.11	0.03	0.1	0.0949
4	0.16	0.02	0.3	0.258
5	0.20	0.01	0.0	0.0249
6	0.32	0.11	0.0	0.0483

Table 3. Factors to calculate S/IVOC emissions from POA emissions from Shrivastava et al. (2011), Mayet al. (2013a, c) and newly calculated factors for this study.

Table 4. Configurations for OA and aerosol activation sensitivity simulations. All simulations except for CB05-SORG-DH contain the VBS treatments for OA.

Name	Gas-	ΔH <sub>vap</sub>	VBS	FF	POA	Aerosol	Cumulus
	Phase				emissions	activation	Scheme
CB05-	CB05	30 kJ mol <sup>-1</sup>	-	-	Original	ARG00	Grell-Freitas
SORG-DH					NEI		
CB05-VBS-	CB05	30 kJ mol <sup>-1</sup>	SOA	-	Original	ARG00	Grell-Freitas
DH					NEI		
CB05-POA-	CB05	30 kJ mol <sup>-1</sup>	SOA/	-	1.5×	ARG00	Grell-Freitas
DH			POA				
CB05-POA	CB05	Epstein et	SOA/	-	1.5×	ARG00	Grell-Freitas
		al. (2010)	POA				
CB05-	CB05	Epstein et	SOA/	50%	1.5×	ARG00	Grell-Freitas
50%FF		al. (2010)	POA				
CB05-	CB05	Epstein et	SOA/	10%	1.5×	ARG00	Grell-Freitas
10%FF		al. (2010)	POA				
CB05-	CB05	Epstein et	SOA/	25%	$1.5 \times$	ARG00	Grell-Freitas
25%FF		al. (2010)	POA				
CB05-	CB05	Epstein et	SOA/	25%	3.0×	ARG00	Grell-Freitas
25%FF-EM3		al. (2010)	POA				
CB6-25%FF-	CB6	Epstein et	SOA/	25%	3.0×	ARG00	Grell-Freitas
EM3		al. (2010)	POA				
SAPRC07-	SAPRC07	Epstein et	SOA/	25%	3.0×	ARG00	Grell-Freitas
25%FF-EM3		al. (2010)	POA				
CB05-	CB05	Epstein et	SOA/	25%	3.0×	FN05	MSKF
25%FF-EM3		al. (2010)	POA				
(FN05)							
CB05-	CB05	Epstein et	SOA/	25%	3.0×	FN05/	MSKF
25%FF-EM3		al. (2010)	POA			BA10	
(FN05/							
BA10)							
CB05-	CB05	Epstein et	SOA/	25%	3.0×	MN14	MSKF
25%FF-EM3		al. (2010)	POA				
(MN14)							

The suffix "DH" in the case names refer to cases with the default  $\Delta H_{vap}$  of 30 kJ mol<sup>-1</sup>, otherwise with the semi-empirical correlation by Epstein et al. (2010). The simulations without the suffix "POA" or "FF" indicate cases with nonvolatile default POA emissions. The suffix "POA" in the case names refer to cases with semivolatile POA. The suffix "FF" in the case names refer to cases with semivolatile POA and with fragmentation and functionalization treatments, and the suffix "EM3" in the case names refer to cases with 3 times the original NEI POA emissions to take into account for missing S/IVOC species. "-" indicates not applicable.

Case	Mean Obs	Mean Sim	Corr	<b>NMB</b> (%)	NME (%)
	0	C against IMP	ROVE		
CB05-SORG-DH	0.88	0.28/ 0.19	0.26	-68.1/ -78.7	73.9/ 80.9
CB05-VBS-DH	0.88	1.19/ 0.79	0.51	34.9/ -10.1	75.5/ 52.3
CB05-POA-DH	0.88	0.89/ 0.59	0.51	1.1/ -32.6	52.4/ 59.0
CB05-POA	0.88	1.05/ 0.70	0.51	18.9/ -20.7	63.2/49.2
CB05-10%FF	0.88	1.05/ 0.70	0.51	19.4/ -20.4	63.0/ 49.1
CB05-25%FF	0.88	0.86/ 0.57	0.49	-2.9/ -35.2	54.6/51.4
CB05-50%FF	0.88	0.56/ 0.37	0.45	-36.4/ -57.6	54.4/ 62.6
CB05-25%FF-EM3	0.88	1.09/ 0.73	0.47	23.8/ -17.5	65.9/ 50.2
CB6-25%FF-EM3	0.88	1.06/ 0.71	0.48	20.5/ -19.6	49.4/ 63.7
SAPRC07-25%FF-EM3	0.88	1.00/ 0.67	0.46	13.3/ -24.4	60.1/ 50.4
	T	C against IMP	ROVE		
CB05-SORG-DH	1.03	0.44/ 0.34	0.30	-57.6/ -66.7	67.9/72.3
CB05-VBS-DH	1.03	1.34/ 0.94	0.52	30.6/ -8.0	70.3/ 51.1
CB05-POA-DH	1.03	1.13/ 0.83	0.52	10.2/ -18.7	58.5/48.7
CB05-POA	1.03	1.29/ 0.94	0.53	25.6/ -8.5	63.8/48.3
CB05-10%FF	1.03	1.29/ 0.94	0.53	25.9/ -8.2	63.8/ 48.2
CB05-25%FF	1.03	1.09/ 0.83	0.51	6.8/ -21.6	55.2/48.2
CB05-50%FF	1.03	0.80/ 0.61	0.47	-22.0/ -40.2	50.8/ 53.4
CB05-25%FF-EM3	1.03	1.32/ 0.97	0.49	29.7/ -5.7	50.7/ 66.9
CB6-25%FF-EM3	1.03	1.30/ 0.95	0.50	27.2/ -7.3	65.2/ 50.0
SAPRC07-25%FF-EM3	1.03	1.23/ 0.90	0.48	20.6/ -11.9	61.4/49.4
		TC against S	ГN		
CB05-SORG-DH	2.71	1.34/ 1.10	0.29	-50.6/ -59.4	60.1/ 64.9
CB05-VBS-DH	2.71	3.35/ 2.44	0.47	23.7/ -5.8	53.1/42.0
CB05-POA-DH	2.71	2.88/ 2.19	0.47	6.2/ -19.0	45.5/41.6
CB05-POA	2.71	3.03/ 2.30	0.46	11.7/ -15.3	44.6/39.9
CB05-10%FF	2.71	3.03/ 2.30	0.46	11.8/ -15.3	44.5/ 39.8
CB05-25%-FF	2.71	2.66/ 2.05	0.44	-1.8/ -24.3	41.5/ 42.0
CB05-50%-FF	2.71	2.07/ 1.65	0.39	-23.8/ -39.1	43.9/ 49.4
CB05-25%FF-EM3	2.71	3.27/ 2.45	0.41	20.5/ -9.5	49.7/ 41.7
CB6-25%FF-EM3	2.71	3.39/ 2.45	0.34	24.9/ -6.4	54.8/45.5
SAPRC07-25%FF-EM3	2.71	3.00/ 2.28	0.41	10.7/ -16.1	45.2/ 42.0

Table 5. Range of statistics for OA/OC ratios of 1.4 and 2.1 (1.4/2.1) for May to June 2010. All simulations use the ARG00 aerosol activation scheme and the Grell-Freitas cumulus parameterization.

Case	Mean Obs	Mean Sim	Corr	<b>NMB</b> (%)	NME (%)			
Bakersfield								
CB05-SORG-DH	0.51	5.9e-04	-0.15	-100	100%			
CB05-VBS-DH	0.51	0.67	0.41	32.5	62.0			
CB05-25%FF-EM3	0.51	0.24	-0.01	-52.0	61.0			
CB6-25%FF-EM3	0.51	0.28	-0.04	-45.8	59.0			
SAPRC07-25%FF-EM3	0.51	0.24	-0.16	-53.1	63.0			
Pasadena								
CB05-SORG-DH	0.63	0.04	-0.07	-94.0	94.0			
CB05-VBS-DH	0.63	0.54	0.09	-14.5	64.3			
CB05-25%FF-EM3	0.63	0.54	-0.2	-14.4	66.2			
CB6-25%FF-EM3	0.63	0.62	-0.2	-2.1	70.0			
SAPRC07-25%FF-EM3	0.63	0.62	0.03	-1.4	70.5			

Table 6. Statistics for evaluation at Bakersfield and Pasadena sites. A bar chart of daily average obs <u>vs.versus</u> sim values can be found in Figure 4.

CASTNET Max 1-h O <sub>3</sub>								
Case	Mean Obs	Mean Sim	Corr	NMB (%)	NME (%)			
CB05-25%FF-EM3	51.8	43.3	0.54	-16.3	21.9			
CB6-25%FF-EM3	51.8	41.9	0.52	-19.1	24.1			
SAPRC07-25%FF-EM3	51.8	48.3	0.51	-6.7	21.1			
	CAS'	TNET Max 8-	h O <sub>3</sub>					
Case	Mean Obs	Mean Sim	Corr	<b>NMB (%)</b>	NME (%)			
CB05-25%FF-EM3	47.4	43.0	0.54	-9.3	18.9			
CB6-25%FF-EM3	47.4	41.8	0.53	-11.8	20.6			
SAPRC07-25%FF-EM3	47.4	47.9	0.50	1.0	19.8			
	Α	QS Max 1-h C	<b>)</b> <sub>3</sub>					
Case	Mean Obs	Mean Sim	Corr	<b>NMB (%)</b>	NME (%)			
CB05-25%FF-EM3	51.0	49.9	0.55	-2.1	18.2			
CB6-25%FF-EM3	51.0	51.5	0.43	1.0	20.8			
SAPRC07-25%FF-EM3	51.0	59.3	0.44	16.4	26.1			
AQS Max 8-h O <sub>3</sub>								
Case	Mean Obs	Mean Sim	Corr	<b>NMB (%)</b>	NME (%)			
CB05-25%FF-EM3	46.2	46.0	0.54	-0.4	18.6			
CB6-25%FF-EM3	46.2	47.4	0.47	2.6	20.3			
SAPRC07-25%FF-EM3	46.2	53.7	0.46	16.3	25.4			
	IN	<b>IPROVE PM</b>	2.5					
Case	Mean Obs	Mean Sim	Corr	<b>NMB (%)</b>	NME (%)			
CB05-25%FF-EM3	4.9	3.8	0.64	-22.0	40.6			
CB6-25%FF-EM3	4.9	4.1	0.65	-16.5	39.6			
SAPRC07-25%FF-EM3	4.9	3.5	0.60	-28.5	42.9			
		STN PM <sub>2.5</sub>		•				
Case	Mean Obs	Mean Sim	Corr	NMB (%)	NME (%)			
CB05-25%FF-EM3	11.1	8.8	0.48	-20.6	40.7			
CB6-25%FF-EM3	11.1	10.0	0.37	-9.3	44.3			
SAPRC07-25%FF-EM3	11.1	7.7	0.40	-30.5	45.2			
AQS PM <sub>10</sub>								
Case	Mean Obs	Mean Sim	Corr	<b>NMB (%)</b>	NME (%)			
CB05-25%FF-EM3	24.6	7.3	0.08	-70.2	73.5			
CB6-25%FF-EM3	24.6	8.0	0.09	-67.7	71.8			
SAPRC07-25%FF-EM3	24.6	6.9	0.09	-71.9	74.8			

Table 7. Statistics for max 1-h and max 8-h  $O_3$  for simulations with different gas-phases against CASTNET and AQS for May to June 2010.

Case	Mean Obs	Mean Sim	Corr	NMB (%)	NME (%)
CB05-SORG-DH	162.1	96.0	0.28	-40.8	50.4
CB05-VBS-DH	162.1	106.0	0.28	-34.6	50.6
CB05-POA-DH	162.1	115.0	0.29	-29.1	47.4
CB05-POA	162.1	117.3	0.29	-27.7	47.3
CB05-10%FF	162.1	117.1	0.29	-27.8	47.2
CB05-25%-FF	162.1	116.4	0.29	-28.2	47.3
CB05-50%-FF	162.1	114.7	0.29	-29.2	47.4
CB05-25%FF-EM3	162.1	116.2	0.29	-28.3	47.3
CB6-25%FF-EM3	162.1	110.4	0.30	-31.9	47.3
SAPRC07-25%FF-	162.1	77.3	0.26	-52.3	55.8
EM3					

Table 8. Statistics for model evaluation for simulated CDNC against MODIS-derived CDNC from Bennartz (2007). All cases use the Grell-Freitas cumulus parameterization.

Table 9. Statistics for simulated CDNC for CB05-25% FF-EM3 against MODIS-derived CDNC from Bennartz (2007) for May to June 2010. <u>All cases use the MSKF cumulus parameterization</u>.

Case	Mean Obs	Mean Sim	Corr	<b>NMB (%)</b>	NME (%)
ARG00	162.1	104.8	0.31	-35.4	49.9
FN05	162.1	173.8	0.26	7.1	93.0
FN05/BA10	162.1	160.8	0.27	-0.8	87.9
MN14	162.1	168.9	0.27	4.2	89.6

# Figures

0.0

CB05\_SORG\_DH

CB05\_VBS\_DH

CB05\_POA\_DH

CB05\_POA



• OA/OC = 1.4 • OA/OC = 2.1 Figure 1. Sim OC and TC concentrations against observations from IMPROVE and STN under two OA/OC ratios: 1.4 and 2.1, resulting in a range of possible OC or TC values denoted by the gray bars. The obs OC or TC is denoted by the horizontal dotted line.

CB05\_FF25%

CB05\_FF50%

CB05\_FF25%\_EM3 CB6\_FF25%\_EM4 SAPRC07\_FF25%\_EM5

CB05\_FF10%



the case CB05\_25%FF\_EM3.



Figure 3. POA/OA ratios of sim data from various sensitivity simulation cases.



Figure 4. Comparison of obs SOA vs. versus sim SOA at CalNex sites in Bakersfield and Pasadena in California.



Figure 5. Time series of OH and diurnal plots of OH and HO2 at Pasadena, CA during CALNEX, 2010.



Figure 6. Spatial plots of several gas and aerosol species for the three cases with different gas-phase mechanisms.



Figure 7. Timeseries plots of IMPROVE OC <u>vs.versus</u> simulated OC at selected sites from sensitivity simulations of different gas-phase mechanisms. The colored bands represent the range of OC values for ratios 1.4 to 2.1.



Figure 8. Timeseries plots of STN TC <del>vs.</del><u>versus</u> simulated TC at selected sites from sensitivity simulations of different gas-phase mechanisms. The colored bands represent the range of OC values for ratios 1.4 to 2.1.



Figure 9. Impact of different VBS cases on CDNC in warm clouds. The plots show the differences between the different sensitivity simulations and CB05\_VBS-DH.



Figure 10. Spatial plots of total column CDNC, total surface OA and total inorganic PM<sub>2.5</sub> concentrations from simulations with different gasphase mechanisms.



Figure 11. Spatial plots for MODIS-derived CDNC from Bennartz (2007) and simulated in-cloud column CDNC from CB05\_25%FF\_EM3 ARG00, FN series, and MN14 from May to June 2010.



Figure 12. Spatial plots of MODIS CCN and AOD against simulated CCN and AOD from MN14 with CB05\_25% FF\_EM3.