

Author response to the reviews of the paper “Simulating warming climate scenarios with intentionally biased bootstrapping and its implications for precipitation”

(Manuscript # gmd-2016-188)

Taesam Lee

Line 53. White space missing

Reply: The space is added accordingly.

Line 58 Clausis-Clapeyron -> Clausius Clapeyron

Reply: The word has been modified accordingly as Clausius-Clapeyron. Note that the author do not remove ‘-’ mark between two name because this relation has been popularly employed as is.

Line 152 $\text{Sigma}^2 = E(X^2) - (EX)^2$

Shouldn't it be $\text{Sigma}^2 = E(X^2) - E(X)^2$? Please check

Reply: The author checked the equation once again and find no error in the equation. It is right with $E(X^2)$.

Figure 3 legend caption. What are the red crosses? Please explain how are defined the boxes, the whiskers, the line in the boxes.

*Reply: The caption has been modified accordingly with adding the sentence as below:
“Boxes indicate the interquartile range (IQR), and whiskers extend to $\pm 1.5IQR$. The horizontal lines inside the boxes depict the median of the data. Data beyond the fences ($\pm 1.5IQR$) are indicated by a plus symbol (+), which represent outliers.”*

Figure 4 legend caption. Please explain how are defined the boxes, the whiskers, the line in the boxes.

*Reply: The caption has been modified accordingly with adding the sentence as below:
Boxes indicate the interquartile range (IQR), and whiskers extend to $\pm 1.5IQR$. The horizontal lines inside the boxes depict the median of the data. Data beyond the fences ($\pm 1.5IQR$) are indicated by a plus symbol (+), which represent outliers.*

Author response to the reviews of the paper “Simulating warming climate scenarios with intentionally biased bootstrapping and its implications for precipitation”

(Manuscript # gmd-2016-188)

Taesam Lee

Reviewer #1

D. Defrance (Referee)

This article presents a statistical method to determine local climate change from global observations. With this approach, the Intentionally Biased Bootstrapping (IBB) and some hypothesis, the author estimates the future temperature and precipitation at a local point. The article is clearly divided into several parts: a good description of the method, the complete procedure to permit to everyone to use easily it and a good application on the South Korea to validate the method with a good description of the results. The methodology is precisely described but some information will permit to improve the comprehension. I suggest to publish this article in GMD with minor revision. The different remarks and suggestions are described below.

Reply: The author appreciates this reviewer’s generous comment. The author tried his best efforts to improve the manuscript. Hope this improvement is satisfactory to this reviewer.

Some questions

Line 31: To specify that the temperature from GCM is relatively accurate as you mention in the conclusion

Line 54: In some places, such as the Sahel, the increasing in temperature results from global warming but also from feedback related to the reduction of precipitations. It is perhaps too generalist to assert that everywhere the increasing in temperature will be followed by an increasing in precipitation with the self-order of magnitude. Can this depend on the type of precipitation or the origin (e.g. monsoon system or stratiform precipitation) ?

Reply: The proposed IBB method does not postulate that the temperature increase means the increase of precipitation. The method employs the empirical relation between temperature and precipitation. When an observed temperature increases and an observed precipitation decreases, the same reverse relation can be reproduced through the proposed IBB method. The author considers that the proposed method is not physical-based method so that the type of precipitation cannot be taken into consideration.

Line 78: In the methodology, some hypothesis must be mentioned: - The method is only based on the temperature mean. If in the future the extremes of temperature increases (warmest and coldest), the method does not take this into consideration. - For the precipitation, the evolution is in relation with only the temperature evolution in the methodology and the meso-scale change is not supported.

The author really appreciates this reviewer’s insightful comment. No physical mechanisms can be included. This limitation was discussed at the conclusion section.

“The proposed IBB method is not a physical-based method but a statistical simulation

approach in which a physical mechanism of precipitation cannot be taken into consideration. Substantial modification might be required to accommodate this mechanism.”

Line 160: for the block bootstrapping technique to simulate the temperature, I would like a better description of the method with one or two sentences because it is easier to read the entire method rather than reading into the references.

Reply: The author totally agrees with this reviewer’s comment. Simple sentences were added accordingly as follows:

“Bootstrapping is a random sampling with replacement and block bootstrapping is to resample blocks. Each block contains a set of predictor and predictand like a regression. Here, temperature and precipitation can be set as a block and they act as predictor and predictand, respectively.”

The author hopes that this modification is satisfactory to this reviewer

Line 191: Data description, you describe the available data (74 locations) and you give 1283 mm a year but you select 54 datasets with a good hypothesis (> 30 years available data). Is the precipitation mean the same with the only 54 datasets? I suggest to insert directly the selected datasets in the beginning of the paragraph with the hypothesis and the annual mean.

Reply: The author appreciate this reviewer’s detailed comment. Official annual mean precipitation of South Korea (1283mm) is announced by KMA, not calculated from the current study. The sentence was modified accordingly as follows:

“In the current study, weather stations that record temperature and precipitation in South Korea (54 locations) and that are managed by the Korea Meteorological Administration (KMA) and whose length is more than 30 years were employed. South Korea is located in Far East Asia and has a mean annual precipitation of 1283 mm from KMA.”

The author hopes that this modification is satisfactory to this reviewer

Line 250: you very accurately write that the test period is relatively short and not enough of high values of annual temperature. Did you tested a longer test period with a short validation period e. g. 20 years test period 1976-1997 and validation period 1998-2008 ?

Reply: The author really appreciates this reviewer’s pinpointing comment. 20 years was also tested with no difference from the current test. 15 years (the test period that has been used in the current study) and 20 years are not much different in analyzing the long-term change.

Line 335: In the conclusion, the limits of the method in terms of variability of extremes should be recreated. This limit associated with IBB can still be disturbing for some applications such as extreme floods. Figure 3 and 4, there are many data on it and it is not easy to analyse it for the reader. Maybe to classify the stations by order of error could permit to better interpret the results. I am not a good example to suggest to you a good representation of the results.

Reply: The author really appreciates this reviewer's insightful comment. The authors consider that long-term variability of hydrological extremes can be derived from the IBB method when it is related with other variables such as precipitation. But no physical mechanisms can be included as this reviewer pointed in the previous comment. This limitation and possible extension were discussed at the conclusion as follows:

“The proposed IBB method is not a physical-based method but a statistical simulation approach in which a physical mechanism of precipitation cannot be taken into consideration. Substantial modification might be required to accommodate this mechanism. Also, a possible extension of the current study must be on analyzing the future variation of hydrological extreme events (e.g. extreme floods). If a long-term variation of hydrological extreme events is related with precipitation, one can derive the variation from the IBB method.”

Hope this reviewer satisfactory to this modification.

Technical notes

Line 58: 1 hour intensity

Reply: It was modified as 'the intensity of hourly precipitation'. Hope this modification is satisfactory to this reviewer.

Line 64: for this paragraph, a reference could be appreciated

Reply: A reference is added accordingly.

Line 98: local linear smoothing (Cai, 2001)

Reply: It was modified as 'local linear regression'.

Line 208: but employed in comparison ? Can you use validation ?

Reply: The author appreciates this reviewer's detailed comment. 'validation' was used now according to this reviewer's comment.

Author response to the reviews of the paper “Simulating warming climate scenarios with intentionally biased bootstrapping and its implications for precipitation”

(Manuscript # gmd-2016-188)

Taesam Lee

Reviewer #2

M. A. Ben Alaya (Referee)

In this paper, the author presents a statistical non parametric resampling approach called intentionally biased bootstrapping (IBB) to simultaneously simulate temperature and precipitation at a single site taking into account the increase of the temperature according to observed global warming data. The manuscript is well organized and the methodology is adequate, reasonable and clearly presented. The problematic and the application are of great interest for GMD. Hence I suggest to publish this paper. However, there are a few statements that don't entirely ring true, and I'd like the author to address these a bit more carefully. Also, drawbacks of the proposed method should be mentioned and discussed.

Reply: The author appreciates this reviewer's generous comment. The author tried his best efforts to improve the manuscript. Hope this improvement is satisfactory to this reviewer.

Below I list relatively minor points that could be addressed with some small revisions to the text and a few more figures:

1- Line 31: “The temperature variable is the most reliable of the GCM outputs”. I'm not sure that this statement is true.

Reply: The author really appreciates this reviewer's detailed comment. The sentence was modified accordingly as:

“The temperature variable is more reliable than other variables in GCM outputs.”

2- Line 57: I agree that moisture availability increases at the same rate with warming through the Clausius-Clapeyron (C-C) relation. Nevertheless this does not guarantee that precipitation intensity should also increase at the same rate, this presumably assumes stationarity of precipitation efficiency.

Reply: The author totally agrees with this reviewer's comment. The sentence was circumvented as follows:

“ From the Clausius-Clapeyron (C-C) relation, saturation vapor pressure increases by 6-7% for each 1°C increase in temperature and rainfall intensity also increase in a similar rate with warming (Trenberth and Shea, 2005).”

3- The proposed approach is based on the assumption that only the mean of observed temperature changes in the future, and assumes a static variance in the future. This assumption should be mentioned. Indeed the proper reproduction of the temporal variability is a very important issue, because a poor representation of the temporal variability could leads to a poor representation of extreme events.

Reply: The author really appreciates this reviewer's insightful comment. The limitation and its possible development is discussed at the conclusion section as the below. Hope this modification is satisfactory to this reviewer.

"The proposed IBB method is conditioned and assumed only on the mean temperature change. A further scheme can be developed to consider changes of multiple variables with classifying the conditions of interested variable."

For the relation of the temporal variability and extreme events, the author consider that this reviewer's comment can be true but not always as far as this reviewer's viewpoint. Further study relates on this issue can be studied.

4- Line 166: "Unlike for the case of temperature, there is no variance reduction in the resampled precipitation data because the precipitation data are not conditionally resampled"; I'm not sure that this statement is true. The existence of dependence between precipitation and temperature which motivates this work implies the existence of a concordance in the ranks of these variables. In the case of dependence there will always be some reduction in the variability of precipitation using the IBB technique. I ask the author to verify this fact by comparing the observed variance and the simulated one in the case of precipitation.

Reply: The author really appreciates this reviewer's insightful comment. The author compared the observed variance with the simulated one for all the 54 stations. No significant variance reduction was observed and even some stations (15 stations) present variance inflation (i.e. simulated variance is bigger than the observed variance). Therefore, the author consider that the statement can be true but with a little less certainty. The sentence was modified as: "not much significant variance reduction is expected in the resampled precipitation data because the precipitation data are not conditionally resampled."

5- The proposed approach is not appropriate to simulate change in extreme events, indeed as it is the case for most resampling approach the IBB technique suffers from the inability to simulate values that are more extreme than those observed.

Reply: The author really appreciates this reviewer's insightful comment. The authors consider that long-term variability of extremes can be derived from the IBB method when it is related with other variables such as precipitation. But it might be limited since no physical mechanisms can be included. This limitation and possible extension were discussed from this reviewer and the other reviewer's comment at the conclusion as follows:

"The proposed IBB method is not a physical-based method but a statistical simulation approach in which a physical mechanism of precipitation cannot be taken into consideration. Substantial modification might be required to accommodate this mechanism. Also, a possible extension of the current study must be on analyzing the future variation of hydrological extreme events (e.g. extreme floods). If a long-term variation of hydrological extreme events is related with precipitation, one can derive the variation from the IBB method."

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Simulating climate warming scenarios with intentionally biased bootstrapping and its implications for precipitation

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29

Abstract

30 The outputs from GCMs provide useful information about the rate and magnitude of future climate
31 change. The temperature variable is ~~morethe most~~ reliable ~~of the GCM outputs~~ than other variables
32 in GCM outputs. However, hydrological variables (e.g., precipitation) from GCM outputs for
33 future climate change possess an uncertainty that is too high for practical use. Therefore, a method,
34 called intentionally biased bootstrapping (IBB), that simulates the increase of the temperature
35 variable by a certain level as ascertained from observed global warming data is proposed. In
36 addition, precipitation data was resampled by employing a block-wise sampling technique
37 associated with the temperature simulation. In summary, a warming temperature scenario is
38 simulated and the corresponding precipitation values whose time indices are the same as the one
39 of the simulated warming temperature scenario. The proposed method was validated with annual
40 precipitation data by truncating the recent years of the record. The proposed model was also
41 employed to assess the future changes in seasonal precipitation in South Korea within a global
42 warming scenario as well as in weekly time scale. The results illustrate that the proposed method
43 is a good alternative for assessing the variation of hydrological variables such as precipitation
44 under the warming condition.

45

46

47 **1. Introduction**

48 The complex influence of human actions on the climate system is well represented through global
49 climate models (GCMs). A number of GCMs demonstrate variations in the large-scale atmospheric
50 circulation and related changes in hydrometeorological variables (Allen and Ingram, 2002; Held
51 and Soden, 2006; Lenderink and Van Meijgaard, 2008). It has been generally accepted that to
52 quantify the range of possible changes in the hydrological cycle (such as precipitation and
53 evaporation) is harder than in temperature (Allen and Ingram, 2002). Furthermore, hydrological
54 variables vary much more in space and time than temperature and difficult to correctly simulate.

55 The relationship between temperature and precipitation has been studied in literature in order
56 to predict the future variations of precipitation under the global warming condition. From the
57 Clausius-Clapeyron (C-C) relation, saturation vapor pressure increases by 6-7% for each 1°C
58 increase in temperature and rainfall intensity ~~should~~ also increases at the same rate in a similar rate
59 with warming (Trenberth and Shea, 2005). Lenderink and Van Meijgaard (2008) presented that
60 the intensity of hourly precipitation exhibit a C-C relation for summer while showing super C-C
61 scaling for winter.

62 These relations are only focused on very short time scale (not more than daily) or generally
63 retrieved from GCM outputs. The behavior of mean precipitation over long-term period such as
64 months and seasons is difficult to predict as temperature increases. It might be beneficial if one
65 could derive the behavior of long-term mean precipitation under warming condition or the range
66 of possible changes (IPCC, 2013).

67 Therefore, a simple method that simulates temperature from observed data is proposed in the
68 current study while increasing temperature up to a certain level as a warming scenario. In addition,

69 precipitation is simulated by employing a block-wise resampling technique (Srinivas and
70 Srinivasan, 2000) associated with the temperature simulation. The resampled covariate,
71 precipitation, forcing the warming condition in a certain level is obtained from the simulation. The
72 proposed approach allows assessing the impact of precipitation as temperature increases with a
73 current climate horizon.

74 The paper is organized as follows. In the next section, the fundamental mathematical
75 background related to bias bootstrapping modeling is presented. The employed data and
76 application methodology are described in section 3. The validation study of the proposed IBB
77 approach is shown in section 4. The results assessing the long-term evolution of seasonal
78 precipitation with simulating weekly temperature and precipitation data are illustrated in section
79 5. Finally, the summary and conclusions are presented in section 6.

80 **2. Methodology**

81 In order to simulate warming scenario, i.e. increasing mean temperature, up to a certain level, the
82 observed data must be sampled with different combination. Intuitively, warmer temperature
83 values are more likely to be resampled among the observations if the mean is increased. Therefore,
84 the proposed method in the current study is to resample the observed data by fixing the mean
85 temperature increment in the resampled dataset by weighting the probability of selection according
86 to its magnitude (see [Figure 1](#)). In addition, the block bootstrapping with precipitation
87 was employed to assess the changes in these variables as temperature increases.

88 **2.1. Intentionally Biased Bootstrapping (IBB)**

89 Bootstrapping (also known as resampling from observed data with replacement) is a statistical
90 method for creating replica datasets from the original data to assess the variability of the quantities

91 of interest without analytical calculation (Davison and Hinkley, 1997; Davison et al., 2003; Ouarda
92 and Ashkar, 1995). This bootstrapping technique has been extended to simulate time series of
93 hydrometeorological variables (Beersma and Buishand, 2003; Lall et al., 1996; Lall and Sharma,
94 1996; Lee and Ouarda, 2011, 2010; Mehrotra and Sharma, 2005). In the current study, the
95 intentionally bias bootstrapping (**IBB**) technique is employed so that the mean of the resampled
96 datasets are varied as needed to simulate a global warming scenario.

97 IBB was proposed by Hall and Presnell (1999) as a class of weighted bootstrapping
98 techniques in order to reduce bias or variance as well as to render some characteristic equal to a
99 predetermined quantity. A good example of IBB is the adjustment of Nadaraya-Watson kernel
100 estimators to make them competitive with local linear [smoothing regression](#) (Cai, 2001). In the
101 current study, IBB was employed to simulate the temperature data from observation by
102 bootstrapping under the constraint of increasing mean value, which indicates warming. The
103 conceptual background of IBB has been employed to simulate future climates of weather analogs
104 (Orlowsky et al., 2010; Orlowsky et al., 2008). In the current study, a IBB method with easy
105 manipulation to simulate increased temperature data is proposed. The mathematical description of
106 the proposed IBB method is the following.

107 Among an n number of observations x_i , where $i=1, \dots, n$, assume resampling the
108 observations with replacement (i.e. bootstrapping) by increasing the mean of the simulated data
109 by as much as Δ_μ ; this implies that higher values have a higher probability of being resampled
110 and lower values have lower selection probability. This IBB can be achieved by assigning different
111 weights $S_{i,n}$ according to the magnitudes of the observations as

$$112 \quad S_{i,n} = i/n \quad (1)$$

113 Note that this assigned weight $S_{i,n}$ plays a role in the selection probability for the observed data in
 114 the IBB procedure after scaling and adjusting it.

115 The mean of the resampled data is

$$116 \quad \tilde{\mu} = \frac{1}{\Psi} \sum_{i=1}^n S_{i,n} x_{(i)} \quad (2)$$

117 where $x_{(i)}$ represents the i^{th} increasing ordered value and $\Psi = \sum_{i=1}^n S_{i,n}$. The amount of the mean
 118 increase δ_{μ} is

$$119 \quad \delta_{\mu} = \tilde{\mu} - \hat{\mu} = \frac{1}{\Psi} \sum_{i=1}^n S_{i,n} x_{(i)} - \frac{1}{n} \sum_{i=1}^n x_i \quad (3)$$

120 To obtain different values of δ_{μ} , the weights can be generalized with the weight order (r) as

$$121 \quad \tilde{\mu}(r) = \frac{1}{\Psi_r} \sum_{i=1}^n S_{i,n}^r x_{(i)} \quad (4)$$

122 where $\Psi_r = \sum_{i=1}^n S_{i,n}^r$. The difference is

$$123 \quad \delta_{\mu}(r) = \tilde{\mu}(r) - \hat{\mu} = \frac{1}{\Psi_r} \sum_{j=1}^n S_{j,n}^r x_{(j)} - \frac{1}{n} \sum_{j=1}^n x_j \quad (5)$$

124 Once the magnitude of the mean increase is given (e.g., temperature increase) as Δ_{μ} , the weight
 125 order ' r ' is estimated accordingly. For example, when the temperature change is obtained from
 126 the GCM outputs and this change is supposed to be propagated into a specific location and a finer

127 time scale, the selection of the weight order can be performed using a meta-heuristic optimization
128 technique with the objective function as

129
$$\text{Minimize } [\Delta_\mu - \delta_\mu(r)]^2 \quad (6)$$

130 In the current study, the harmony search (HS) was used for the meta-heuristic optimization. The
131 performance of the HS in hydrological applications is well reviewed in the literature (Geem et al.,
132 2001; Lee and Geem, 2005, 2004; Lee and Jeong, 2014a; Mahdavi et al., 2007; Yoon et al., 2013a).
133 Note that if $r > 0$, then $\delta_\mu(r) > 0$, which implies a global warming scenario; if $r < 0$, then
134 $\delta_\mu(r) < 0$, which implies a global cooling scenario. When $r < 0$, lower values are resampled more
135 frequently than are higher values, causing the mean of the resampled data to decrease. Furthermore,
136 if r goes to infinity then the maximum of the observations is always selected, and if r goes to
137 negative infinity, only the minimum is chosen.

138 In the IBB procedure, the adjusted scaled weight $\eta_i = S_{i,n}^r / \Psi_r$ is the probability that each i^{th}
139 data point is subject to be selected. In the case of $n=30$, the weights for $i=1, \dots, n$ are shown in
140 [Figure 2](#) with the weight order of $r=0.5$. The figure presents that the probability of being
141 selected (i.e., η_i) is between approximately 0.01 for the lowest values and 0.05 for the highest
142 order values of approximately 0.05 to lead positive bias in the resampled data (e.g., 1.0°C increase).
143 For example, if the number of the simulation is 100 and $\eta_i=0.05$, then the data point will be
144 selected 5 times. A different probability implies a different number of selection for each data point.
145 Subsequently, a different number of selections may lead to variation changes, called variance
146 reduction or inflation. This issue is dealt with in the following section.

147 **2.2. Variance reduction and inflation**

148 Because of the biased selection of higher values, the variance of the resampled data results is
149 reduced (Lee and Jeong, 2014a; Lee and Ouarda, 2010; Lee et al., 2010a; Salas and Lee, 2010;
150 Sharif and Burn, 2006). The estimated variance of the simulated data with IBB is

151
$$\tilde{\sigma}^2(r) = \sum_{j=1}^n \frac{S_{j,n}^r}{\Psi_r} x_{(j)}^2 - \tilde{\mu}^2 \quad (7)$$

152 Note that the variance in Eq. (7) is based on $\sigma^2 = E(X^2) - (EX)^2$. The difference of the variance
153 is

154
$$\delta_{\sigma^2}(r) = \hat{\sigma}^2 - \tilde{\sigma}^2(r) \quad (8)$$

155 where $\hat{\sigma}^2$ is the sample variance of the observed data. To overcome the reduction of the variance
156 in IBB, a random perturbation can be applied to the resampled data X_R as

157
$$X_{R^*} = X_R + \sqrt{\delta_{\sigma^2}(r)} \varepsilon \quad (9)$$

158 where ε is a random variable with a normal distribution $N(0,1)$. Subsequently, the mean and
159 variance of the perturbed data are

160
$$\hat{\mu}_{R^*} = \tilde{\mu} \quad (10)$$

161
$$\hat{\sigma}_{R^*}^2 = \tilde{\sigma}^2 + \delta_{\sigma^2}(r) = \tilde{\sigma}^2 + \hat{\sigma}^2 - \tilde{\sigma}^2(r) = \hat{\sigma}^2 \quad (11)$$

162 **2.3. Block bootstrapping**

163 Bootstrapping is a random sampling with replacement and block bootstrapping is to resample
164 blocks. Each block contains a set of predictor and predictand like a regression. Here, temperature

165 and precipitation can be set as a block and they act as predictor and predictand,
166 respectively.resample random samples

167 When the temperature presumably increases by a certain degree, it is interesting to note how
168 the other weather variables vary. For example, if the temperature is increased by 1°C, the greatest
169 concern in climate research will be how the precipitation will change.

170 To address this question, the block bootstrapping technique for the precipitation variable is
171 adapted (Carlstein et al., 1998; Lee et al., 2010b). Once the temperature is resampled from the
172 observed data at certain times using IBB, the observed precipitation data from the same time are
173 considered (see Figure 2). Unlike for the case of temperature, ~~there is no variance reduction in~~ not
174 much significant variance reduction is expected in the resampled precipitation data because the
175 precipitation data are not conditionally resampled. This block bootstrapping technique is popularly
176 employed in multivariate weather simulations (Lee and Jeong, 2014b; Lee et al., 2012).:-

177 **2.4. Overall Simulation Procedure**

178 The overall simulation procedure of temperature and precipitation data is described in this section.
179 Simple schematic presentation of the procedure is shown in ~~Figure 1~~ Figure 4.

180 Let x_i, y_i ($i=1, \dots, n$) be the observed temperature and precipitation data, respectively. Suppose that
181 the simulation length is the same as the record length (i.e. n) and 100 series need to be simulated.

182 (a) Assume that the increased overall temperature mean is known as Δ_μ .

183 (b) Estimate the weight order (r) from meta-heuristic algorithm (here, Harmony Search) with
184 the objective function of Eq.(6) from the observed temperature data.

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185 (c) Resample the temperature data from the observations with the probability of $S_{i,m}^r$ for i^{th}
186 largest data ($i=1, \dots, n$).

187 (d) Assume that k^{th} largest temperature data $x_{(k)}$ is resampled from step (3) and its
188 corresponding time index of (k) is ' j '. Note that (k) indicates the k^{th} largest value and j
189 indicates the j^{th} time-index value. Then, j^{th} precipitation data, y_j , is resampled
190 simultaneously.

191 (e) Apply Eq.(9) to the resampled temperature data from step(3) (say, $x_{(k)} + \sqrt{\delta_{\sigma^2}(r)}\varepsilon$), if the
192 variance inflation is chosen.

193 Note that the current procedure is explained for the case of no seasonal variability due to
194 simplicity. In other words, the explained procedure above must be applied at each week or each
195 month for weekly or monthly data. The detailed description of the proposed method for the case
196 of monthly precipitation data with the full record is provided in the supplementary material
197 (Supplement A).

198 3. Data description and application methodology

199 In the current study, [54](#) weather stations that record temperature and precipitation in South
200 Korea (~~74 54 locations~~) [with more than 30 years of record length](#) and that are managed by the
201 Korea Meteorological Administration (KMA) ~~and whose length is more than 30 years~~ were
202 employed. South Korea is located in Far East Asia and has a mean annual precipitation of 1283
203 mm [from KMA](#). This country is climatologically influenced by the Siberian air mass during winter
204 and the Maritime Pacific High during summer. Most of the annual precipitation in South Korea
205 falls during the rainy season from June to September due to the occurrence of tropical cyclones,

206 extratropical cyclones, fronts and other weather systems. Because the orographic area in South
207 Korea is heterogeneous and large, the rainfall in South Korea has large spatial and temporal
208 variability (Park et al., 2007; Yoon et al., 2013b). The water resource control system, including
209 climate change, is an important aspect of this study due to the seasonal and spatial variability of
210 rainfall in this country.

211 Datasets shorter than 30 years of data were excluded, after which a total of 54 datasets were
212 employed. The data were extracted from the KMA website (<http://www.kma.go.kr/>). Most of the
213 time spans are approximately 33 years, from 1976 to 2008.

214 The validation study was performed with annual dataset to present the performance of the
215 proposed model with truncating recent years as 1994-2008. The truncated data was not used in
216 simulation but employed in comparison validation. Also, a case study was applied with the weekly
217 dataset of the 54 stations in South Korea. In the application study of the proposed IBB procedure
218 in section 5, (1) 0.5°C and 1.0°C increases in the mean weekly temperature were assumed; (2)
219 weekly temperature datasets were simulated using the assumed temperature increase; (3) weekly
220 precipitation datasets were also simulated along with the weekly temperature dataset as a block.
221 Note that the simulation does include not a gradual change, such as a trend, but the overall mean
222 change. We simulated the weekly time scale so that the data spanned a long enough period to
223 provide a summary of weather statistics and a short enough period to reflect the temporal
224 variability. Furthermore, the observed weekly datasets of temperature and precipitation were
225 aggregated into seasonal time scale data, and the aggregated seasonal data were used to present
226 the seasonal variations in precipitation as temperature increases.

227 Note that although we simulated the temperature with a specific condition of increase (e.g.
228 +0.5°C or +1.0°C), no such restriction was placed on the precipitation, allowing one to determine
229 whether there is any change in precipitation with the condition of increasing temperature. One
230 hundred series were simulated with the same time span as the observations.

231 **4. Validating IBB model with annual data**

232 To further obtain the credibility of the proposed IBB model, we validated the model with truncating
233 the last 15 years (1994-2008) of the annual mean temperature and precipitation data over South
234 Korea. The last truncated 15 years were set as the validation period while the rest of the preceding
235 years as the test period. The dataset of the test period was employed in simulation while the dataset
236 of the validation period is only used in comparison to check how much the proposed model
237 performs. Among others, annual scale data is employed to easily illustrate the performance of the
238 proposed IBB model. At first, some mathematical terms need to be defined to explain the
239 validation procedure as follows.

$$240 \quad D\mu_p^{obs} = \mu\varphi_V - \mu\varphi_T \quad (12)$$

$$241 \quad D\mu_p^{IBB} = \mu\varphi_{IBB} - \mu\varphi_T \quad (13)$$

242 where $\mu\varphi_V$ and $\mu\varphi_T$ are the mean annual precipitation over the validation years and over the test
243 period, respectively, while $\mu\varphi_{IBB}$ is the annual mean precipitation of the IBB simulated data with
244 the record length of the validation years. The same denotation as the precipitation variable is taken
245 for the temperature variable as μT_V , μT_T , μT_{IBB} , $D\mu_T^{obs}$, and $D\mu_T^{IBB}$.

246 The validation procedure is (1) to truncate the 15 years (1994-2008) of annual temperature
247 and precipitation for each station; (2) to estimate the mean differences of the annual temperature
248 and precipitation between the validation period (1994-2008) and the test period (1976-1993),
249 $D\mu_T^{obs}$ and $D\mu_p^{obs}$, respectively; (3) to perform the IBB simulation with the annual precipitation and
250 temperature of the test period conditioned on the estimated mean differences of the temperature
251 between two periods (i.e. $D\mu_T^{obs}$) for each station; and (4) to compare the estimated mean
252 differences of the observed precipitation (i.e. $D\mu_p^{obs}$) with the mean differences between the IBB
253 simulated precipitation and the precipitation for the test period (i.e. $D\mu_p^{IBB}$).

254 The annual mean temperature differences between the validation period and the test period
255 at each station is presented in [Figure 3](#) for the IBB simulated data ($D\mu_T^{IBB}$, boxplot) and
256 the observed data ($D\mu_T^{obs}$, circle). The figure indicates that the IBB model fairly well simulates the
257 temperature data as much as it was intended, except few stations that shows high increase
258 especially with more than one-degree increase (e.g. stations 6 and 7). Note that the employed test
259 period is relatively short and not enough number of high values of annual temperature is included
260 during the test period and this might result the underestimation of the intended temperature
261 increase.

262 In [Figure 4](#), the annual mean precipitation of the observation over the validation
263 period (μp_V , filled blue circle) and the test period (μp_T , filled red triangle) as well as the IBB
264 simulation (μp_{IBB} , boxplot) is illustrated. The result indicates that the observed mean precipitation
265 over the validation period (μp_V) presents higher than the mean for the test period (μp_T) in most

266 of the stations. The IBB simulated data reflects this tendency showing higher mean precipitation
267 than the mean precipitation of the test period though its magnitude shows some difference.

268 The mean of the observed annual precipitation for the validation period at each station and
269 the mean of one hundred IBB simulated data is presented in [Figure 5](#). The top panel
270 presents that the simulated data fairly well reproduce the observed mean of annual precipitation
271 for the validation period (1994-2008). The observed mean difference ($D\mu_p^{obs}$) of the annual
272 precipitation between the test period (1976-1993) and the validation period shown at the bottom
273 panel of [Figure 5](#) fairly matches with the one of the IBB simulated data ($D\mu_p^{IBB}$). Rather
274 high variability at the difference is inevitable due to relatively small record length for both the test
275 period and the validation period. Overall, the validation study implicates that the proposed IBB
276 approach can simulate the future evolution of annual precipitation over South Korea.

277 In [Figure 6](#), the spatial distribution of the differences for the annual mean
278 precipitation is presented with the observed data (i.e. $D\mu_p^{obs}$) and with the IBB simulated data
279 ($D\mu_p^{IBB}$). High increase of annual mean precipitation in the north and south part of the country and
280 small increase and slight decrease in the south part shown in the observed data (left panel) is well
281 reflected in the IBB simulated data (right panel) except that the increase is shown from the IBB
282 simulated data (right panel) in the left south part of the country is not shown in the observed data.
283 Overall, the figure indicates that the spatial pattern of the annual mean precipitation difference
284 from the observed data (see the left panel) is similar to the one from the IBB simulated data (see
285 the right panel).

286 **5. Precipitation changes according to assumed temperature increase**

287 [Figure 7](#) shows the results of the fitted IBB model for the Buan station, located at 35° 44'
288 N and 126° 43' E. The top panel ([Figure 7\(a\)](#)) shows the estimated weight order of each
289 week for the mean temperature data employing the HS meta-heuristic algorithm with the objective
290 function of Eq. (6) while assuming a 0.5°C increase. The estimated values range from 0.2 to 1.3.
291 The mean and standard deviation of the observed and theoretical results (see Eqs. (2) and (7)) with
292 a 0.5°C mean increase are shown in [Figure 7\(b\)](#) and (c), respectively. The predominant
293 annual cycle of the mean weekly temperature is seen in the mean statistics, as shown in [Figure](#)
294 [7\(b\)](#), while the annual cycle of the standard deviation (equivalent to the square root of
295 variance) is not as prominent as the annual cycle of the mean (see [Figure 7\(c\)](#)). Note that
296 the weight order and the standard deviation (see [Figure 7\(a\)](#) and (c)) are highly negatively
297 correlated. In other words, when the standard deviation is small (e.g., at approximately the 23rd
298 week), the weight order is high and vice versa. This result is intuitive in that if the variance is great,
299 the corresponding temperature values differ greatly from each other. Subsequently, the weights of
300 the large values to be selected are not necessarily much different from the weights of the low values
301 in such a case, which induces a low weight order. In [Figure 7\(c\)](#), the variance difference
302 between the observed and theoretical data, as defined in Eq. (8), is shown with a dotted line. This
303 variance difference is inflated to the resampled data, as in Eq. (9). This inflation procedure is
304 optional in assessing the overall trend of annual mean precipitation data regarding climate warming
305 scenarios. However, it might be helpful when the purpose of the study is to evaluate an overall
306 variation of extreme precipitation statistics.

307 The statistics of the simulated data from IBB with the condition of a 0.5°C degree mean
308 temperature increase are shown as a boxplot in [Figure 8](#); the statistics of the observed data
309 are shown in the same figure with dotted lines and cross marks. The mean increases by exactly
310 0.5°C, as intended, and the standard deviation (square root of variance) is well preserved through
311 the variance inflation process (see Eq. (8)). The minima and maxima of the mean weekly
312 temperatures are increased.

313 Shown in [Figure 9](#)(a) are the mean differences between the simulated and observed
314 weekly precipitation with the conditions of 0.5°C and 1.0°C increases at the Buan station. The
315 differences are not significant at the 5% level. However, the mean differences are continuously
316 positive from the 30th to 40th week, which is during the summer season. This result indicates that
317 a seasonal effect on the precipitation change must exist. Therefore, we also extended our study to
318 a seasonal time scale. The mean precipitation differences of all 54 stations are shown for 0.5°C
319 and 1.0°C increases in [Figure 9](#)(b) and (c), respectively. Both plots show a decrease in
320 autumn and increases in the other seasons.

321 For a 1.0°C temperature increase, 61%, 24%, and 45% of the employed stations show a
322 significant increase in mean precipitation for the winter, spring, and summer seasons, respectively.
323 In contrast, the mean temperature decreases during the autumn season. Approximately 30% of the
324 stations experience a significant change in the mean precipitation at the 5% level given a 1.0°C
325 temperature increase. The detailed information is provided in Table 1.

326 The spatial distribution of seasonal mean precipitation differences is presented in [Figure](#)
327 [10Figure 10](#) given the condition of a 1°C temperature increase. An increasing pattern of
328 precipitation during winter (see [Figure 10Figure 10\(a\)](#)) can be seen over the South Korea peninsula.
329 Notably, the eastern and southern coastal areas undergo a significant increase with a 95%
330 confidence interval (± 5.38). Note that the significance interval at each station is different because
331 the variances between stations are different. The detailed significance interval for each station is
332 provided in Table 2. During spring (see [Figure 10Figure 10\(b\)](#)), the northern part of the country
333 shows an increasing pattern while the southwestern and southeastern parts show decreasing
334 patterns, but their magnitudes are not significant (± 15.04). The summer precipitation (see [Figure](#)
335 [10Figure 10\(c\)](#)) undergoes a significant increase in the southwest area of the country (± 29.94). In
336 contrast to the other seasons, a significant decrease in mean precipitation occurs during autumn
337 (see [Figure 10Figure 10\(d\)](#)) throughout the country, especially over the eastern coastal area. The
338 same spatial pattern of seasonal mean precipitation can be observed given the condition of a 0.5°C
339 temperature increase, as in the case of a 1.0°C temperature increase, with little significant change
340 (see [Figure 11Figure 11](#)).

341 The spatial distributions of seasonal precipitation changes seem to be related to the flow
342 direction of the seasonal air mass. In South Korea, winter is influenced primarily by the Siberian
343 air mass with prevailing northwesterly winds, while summer is hot and humid with southeasterly
344 winds.

345 **6. Summary and Conclusions**

346 A simple method is proposed (1) to simulate precipitation given the condition of a mean
347 temperature increase derived from the observations and (2) to address the problem of how the
348 precipitation vary while the temperature is increased through global warming. The results
349 illustrated that a simple IBB technique for the temperature variable incorporating block sampling
350 of precipitation can achieve this objective.

351 The presented technique is valuable because hydrometeorological variables such as
352 precipitation and discharge are difficult to model with current GCMs, while the temperature
353 prediction is relatively accurate. The proposed method can be extended to other
354 hydrometeorological variables as well as other applications, including studies at the global scale.
355 The limit of the proposed method is that the temperature increase is limited since employed data
356 is observational. One possibility for allowing a greater temperature increase than that from the
357 observations is to include neighboring, similar stations or seasons. The author believes that the
358 proposed model can be a good surrogate or competitor in GCM-based climate change impact
359 assessments of hydrometeorological variables.

360 The proposed IBB method is not a physical-based method but a statistical simulation
361 approach in which a physical mechanism of precipitation cannot be taken into consideration.
362 Substantial modification might be required to accommodate this mechanism. The proposed IBB
363 method is conditioned and assumed only on the mean temperature change. A further scheme can
364 be developed to consider the changes of multiple variables with classifying the conditions of
365 interested variable. Also, aAnother possible extension of the current study must be on analyzing
366 the future variation of hydrological extreme events (e.g. extreme floods). ~~When~~ a long-term

367 variation of hydrological extreme events is related with precipitation, the proposed IBB method
368 one can can be used to derive the variation from the IBB method.

369 **7. Code and Data Availability**

370 All the employed code can be provided upon the request to the author of the current study. The
371 employed precipitation and temperature data over South Korea can be downloaded from the KMA
372 website http://www.kma.go.kr/weather/climate/past_cal.jsp .

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376 provided upon the request to the author of the current study.

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462 Table 1. Mean precipitation difference of the observed and simulated data for seasonal data over
 463 all the employed stations in South Korea in case of +1.0 °C mean temperature increase.

Station	Mean Diff				Station	Mean Diff			
	Winter	Spring	Summer	Autumn		Winter	Spring	Summer	Autumn
1	11.2	14.3	20.2	-12.0	28	2.6	12.1	9.6	-4.1
2	3.2	22.4	4.5	0.0	29	4.4	20.6	50.4	-3.8
3	11.0	5.0	21.5	-17.2	30	5.5	11.7	30.0	-4.2
4	1.6	15.7	38.1	-2.3	31	4.4	19.2	15.8	-4.7
5	1.5	11.9	3.9	-6.2	32	4.2	15.9	18.0	-2.0
6	1.7	10.1	28.5	-2.0	33	6.6	16.4	46.1	-4.2
7	1.7	8.2	16.8	-2.3	34	9.5	9.5	32.6	-7.1
8	3.2	22.3	33.6	-3.1	35	6.4	1.7	44.1	-6.8
9	2.3	19.1	15.0	-4.9	36	5.1	-4.2	52.1	-9.4
10	9.8	6.7	21.4	-16.3	37	5.6	7.4	39.9	-9.4
11	2.8	18.8	30.3	-3.3	38	9.2	-4.3	53.8	-3.1
12	5.3	10.8	32.9	-7.2	39	9.6	-3.2	65.0	-5.6
13	5.1	3.5	21.5	-9.3	40	11.5	-9.9	82.2	-6.5
14	9.8	1.2	28.8	-4.5	41	9.1	4.2	33.3	-7.4
15	6.6	-0.9	11.5	-5.1	42	9.6	-11.5	61.2	-8.1
16	5.9	-1.0	32.6	-7.5	43	4.2	12.9	42.7	-3.0
17	10.2	-9.3	26.7	0.6	44	6.3	20.2	33.8	-2.6
18	8.2	-1.7	50.2	-4.5	45	12.9	8.8	10.5	-7.9
19	13.2	-2.7	23.4	0.8	46	5.8	11.2	19.4	-3.8
20	9.8	-4.3	33.1	-0.7	47	3.1	14.3	56.3	-7.0
21	8.1	-15.4	12.4	-4.5	48	7.1	-2.4	14.8	-4.7
22	7.8	-6.0	52.3	-2.3	49	9.0	3.4	68.4	-5.9
23	11.4	-17.5	19.7	-12.6	50	4.2	2.1	31.6	-2.3
24	1.9	11.2	21.1	0.1	51	8.9	5.5	39.5	-3.2
25	2.3	8.6	21.8	-2.4	52	8.6	8.0	78.2	-1.5
26	2.3	8.8	13.4	0.8	53	16.4	6.0	28.8	-4.1
27	2.5	9.3	26.0	-2.9	54	10.5	20.9	23.2	1.7
	Mean confidence interval					±5.38	±15.04	±29.94	±7.01
	# of Significant Stations					33	13	25	16
	(percent)					(61%)	(24%)	(46%)	(30%)

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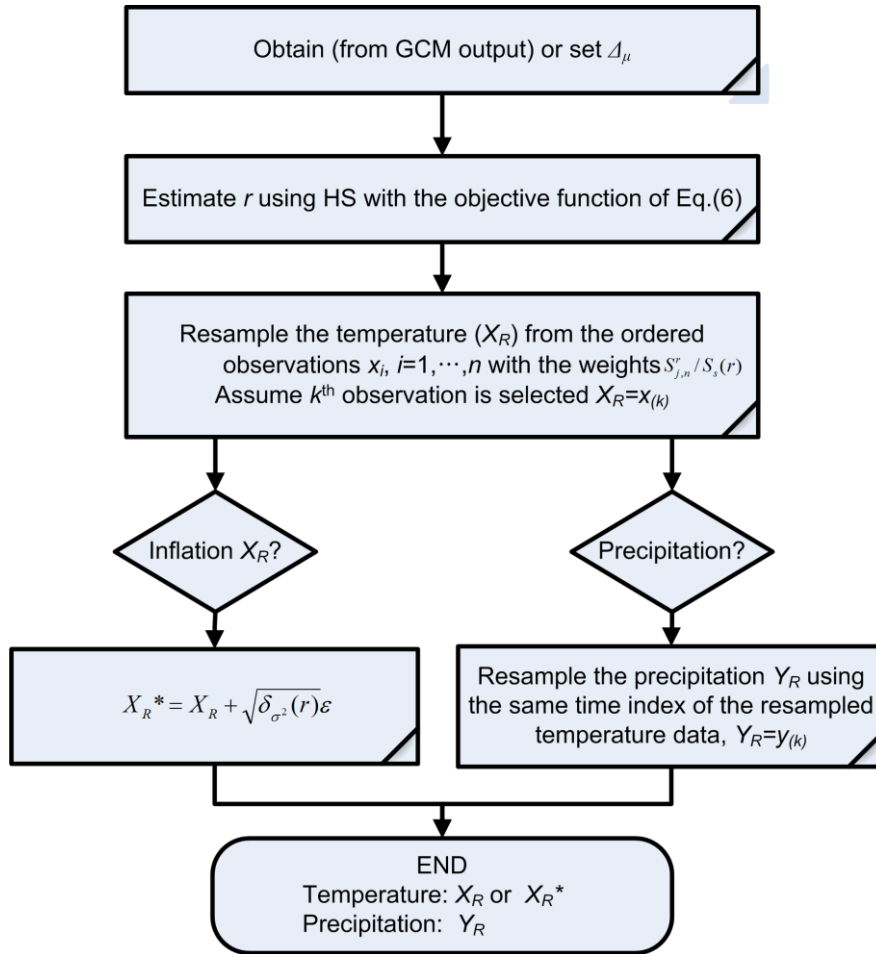
468 Table 2. Confidence interval for mean precipitation difference of the observed and simulated data
 469 for seasonal data.

Station	Winter	Spring	Summer	Autumn	Station	Winter	Spring	Summer	Autumn
1	10.7	12.4	28.4	13.6	28	3.89	14.15	32.45	6.08
2	3.7	13.2	29.0	5.1	29	4.71	14.76	31.49	6.34
3	12.7	10.3	29.6	14.2	30	5.24	14.79	30.39	5.55
4	3.7	14.6	34.7	7.6	31	4.08	14.26	27.61	7.45
5	3.6	12.0	25.9	7.8	32	4.25	14.31	28.31	7.09
6	4.0	12.0	25.3	5.6	33	5.00	15.87	31.29	8.08
7	3.6	14.0	25.9	7.7	34	5.62	13.73	25.75	6.06
8	4.1	13.7	26.4	6.4	35	4.86	12.44	30.64	6.93
9	4.1	14.8	27.1	8.6	36	5.61	12.53	27.52	7.52
10	8.9	10.5	26.7	11.4	37	5.32	12.89	26.21	7.28
11	4.8	14.5	23.0	7.0	38	5.12	13.53	32.37	5.46
12	5.5	15.2	30.7	6.4	39	5.15	15.64	34.46	6.45
13	4.6	13.1	24.6	5.2	40	5.27	20.28	37.15	6.87
14	8.2	12.9	30.9	6.7	41	4.80	20.76	29.50	5.57
15	4.8	12.1	23.6	4.5	42	5.20	21.00	35.75	7.88
16	5.6	12.5	26.9	6.3	43	4.45	15.73	26.47	6.16
17	7.2	15.7	30.1	6.9	44	5.23	14.63	26.25	5.11
18	5.2	15.4	31.9	5.7	45	8.23	11.25	24.05	7.16
19	6.9	20.1	35.1	8.7	46	4.30	10.81	24.10	4.29
20	6.0	19.3	34.3	7.5	47	4.60	11.30	25.36	4.91
21	4.6	15.7	26.5	6.1	48	4.80	11.24	23.40	4.32
22	5.0	19.5	30.1	6.9	49	5.81	12.41	34.88	5.73
23	5.4	22.6	39.4	8.4	50	5.38	14.71	33.37	5.54
24	3.6	17.3	27.5	8.3	51	4.73	15.29	30.09	6.00
25	3.6	13.1	30.8	6.6	52	6.32	17.35	41.62	7.15
26	4.0	13.5	28.2	6.9	53	7.70	29.41	44.00	11.16
27	3.3	13.5	27.7	4.6	54	7.56	23.95	42.12	9.89

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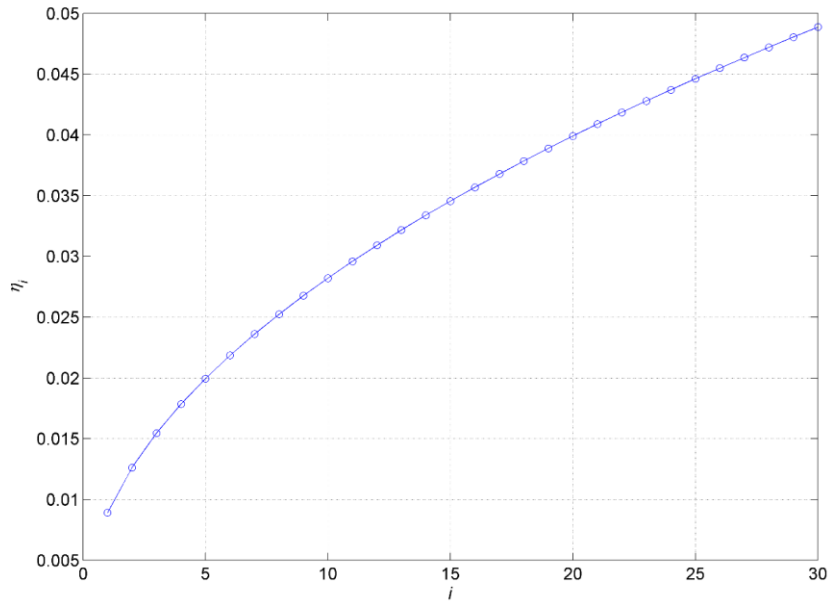
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474 Figure 1. Procedure for the proposed simulation IBB method of temperature and precipitation data.

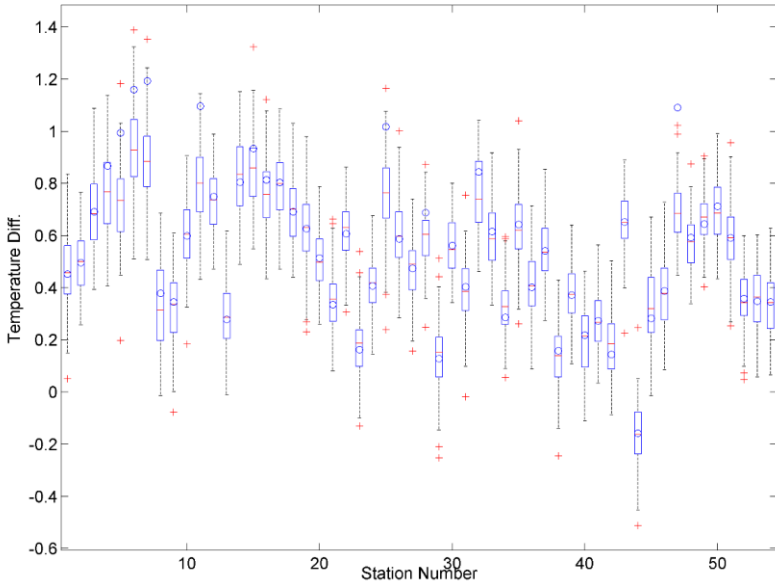
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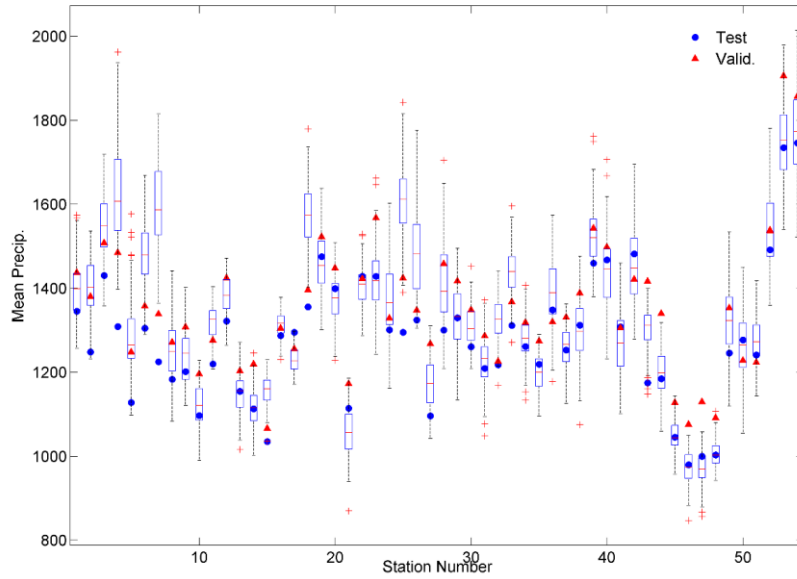
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477 Figure 2. Example of the adjusted scaled weights (η_i) vs. order numbers in the case of $n=30$ and
 478 order weight $r=0.5$. Note that η_i is the probability of being selected and increases as the order is
 479 increased, so that higher values are subject to being selected more often than are lower values,
 480 leading to a positive bias.

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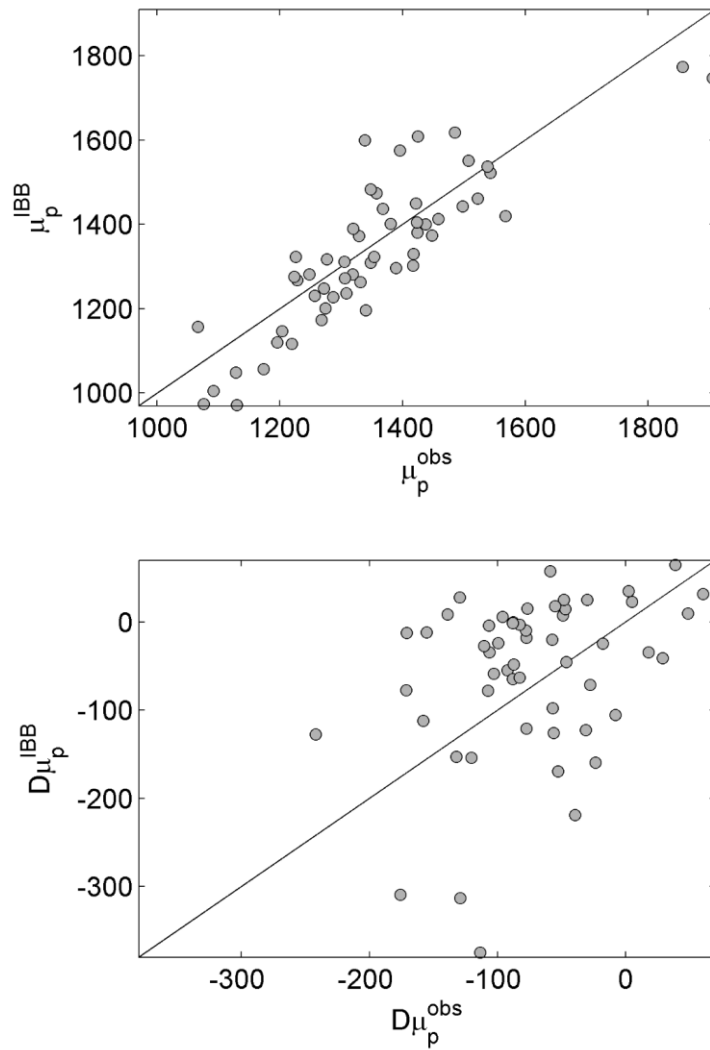
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 483 Figure 3. Annual mean temperature difference between the validation period (1994-2008) and
 484 the test period (1976-1993) for each station for the IBB simulated data (boxplot) and the
 485 observed data (circle). Boxes indicate the interquartile range (IQR), and whiskers extend to +/-
 486 1.5IQR. The horizontal lines inside the boxes depict the median of the data. Data beyond the
 487 fences (+/-1.5IQR) are indicated by a plus symbol (+), which represent outliers.



488

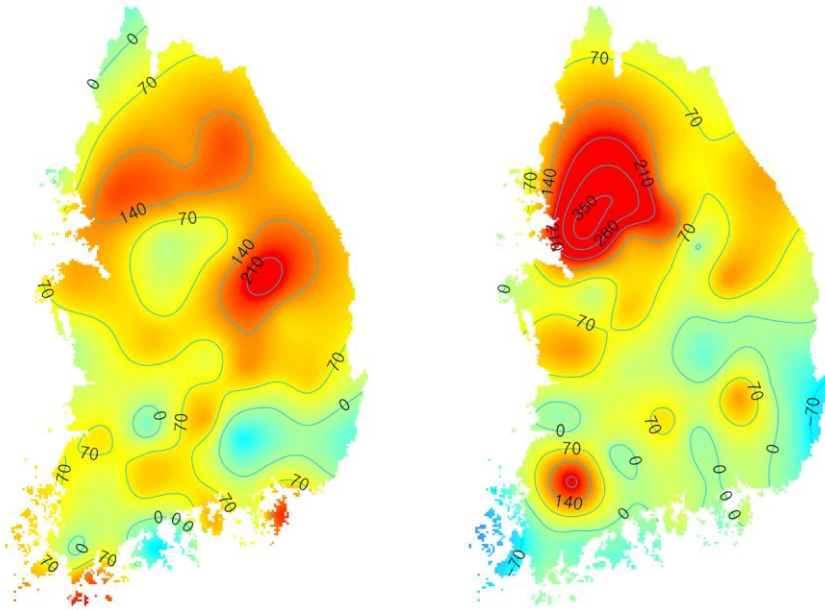
489 Figure 4. Annual mean precipitation of the IBB simulation (boxplot) and the observation over the
 490 validation period (filled blue circle) as well as the test period (filled red triangle) conditioned with
 491 the temperature change (see [Figure 3](#)). Note that the observed mean precipitation over the
 492 validation period (1994-2008) (see the red triangles) shows mostly higher than the mean over the
 493 test period (1976-1993) (see the blue circles). Also, the IBB simulated precipitation (boxplot)
 494 reflects this tendency showing higher than the mean precipitation of the test period (blue circles).
 495 Boxes indicate the interquartile range (IQR), and whiskers extend to ± 1.5 IQR. The horizontal
 496 lines inside the boxes depict the median of the data. Data beyond the fences (± 1.5 IQR) are
 497 indicated by a plus symbol (+), which represent outliers.

498



500 Figure 5. Annual mean precipitation (top panel) during the validation period (1994-2008) and its
 501 difference (bottom panel) with the test period (1976-1993) for the observed data (abscissa) and
 502 the IBB simulated data (ordinate) over all the employed stations in South Korea. For more details
 503 about the difference at the bottom panel, see Eqs. (12) and (13).
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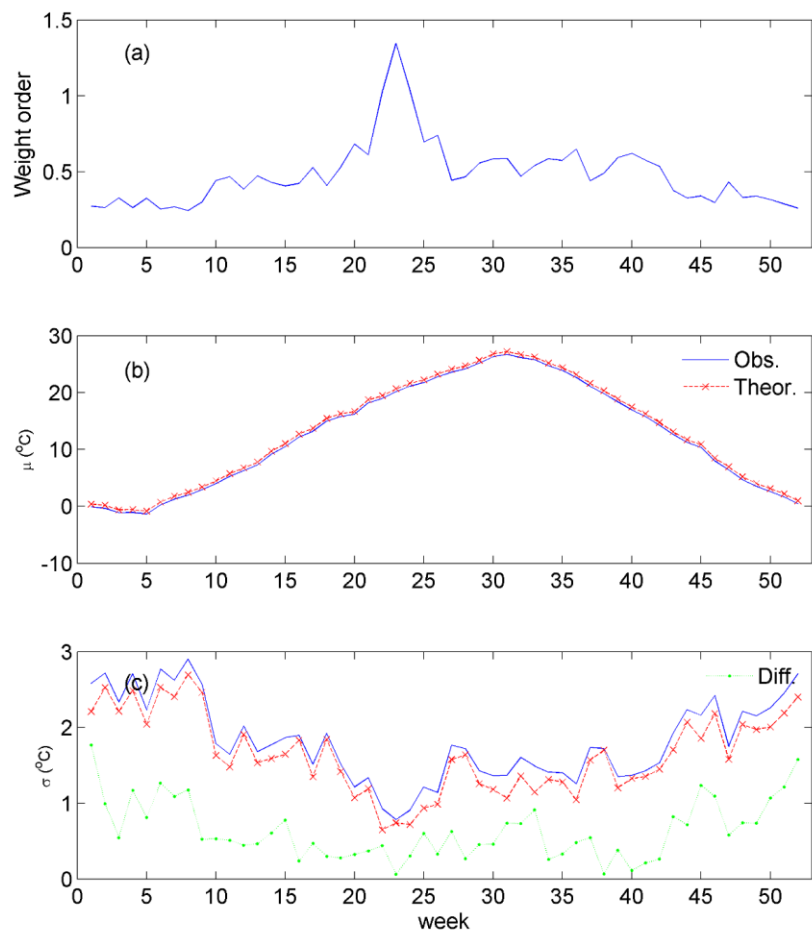


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507 Figure 6. Spatial distributions of annual mean precipitation difference between the validation
508 period (1994-2008) and the test period (1976-1993) for the observed data (left panel) and
509 IBB simulated data (right panel).

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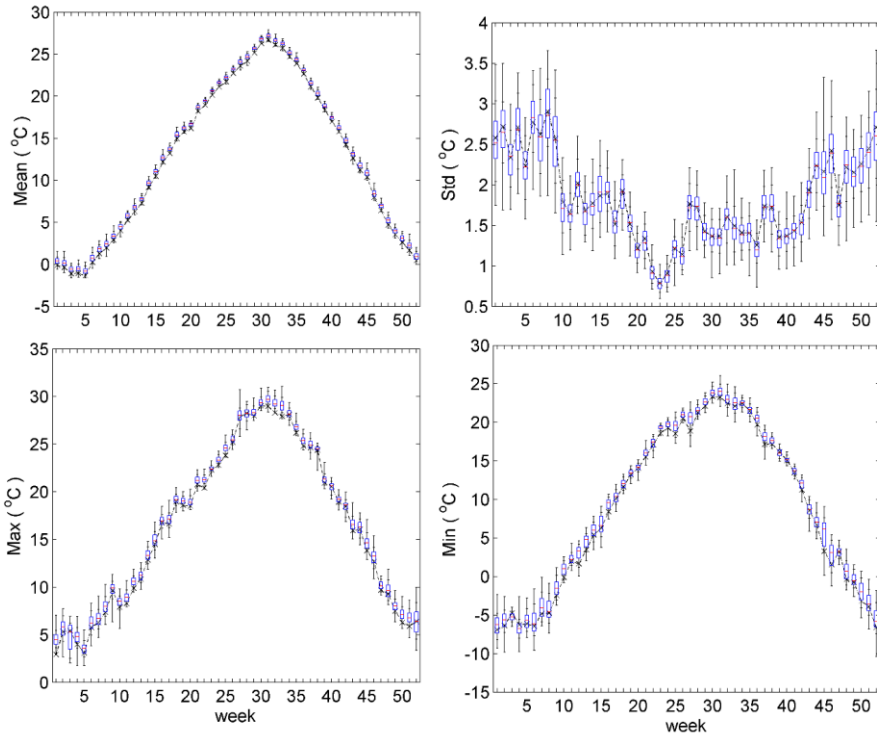
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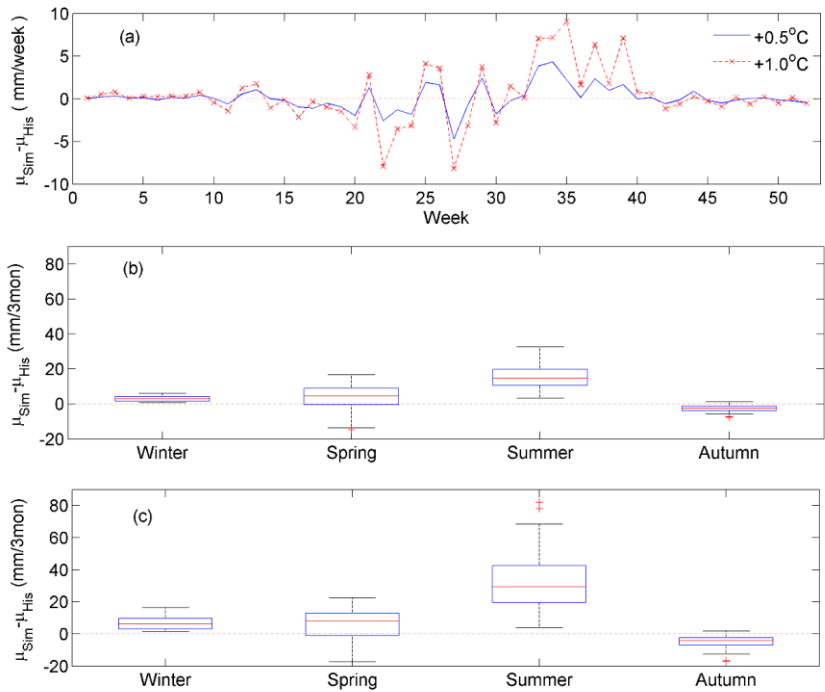
513 Figure 7. (a) Estimated weight order from HS and weekly statistics of (b) mean and (c) variance
 514 for the observed temperature data (solid line) and the theoretical statistics (dashed line with cross)
 515 using Eqs. (2) and (7) for Buan station. The weekly difference in variance between observation
 516 and theoretical (see Eq. (8)) is shown in panel (c) by a dotted line.

517



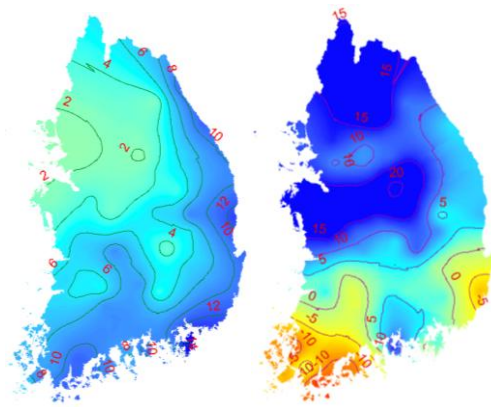
518
 519 Figure 8. The statistics of the observed (dotted line with cross) and generated (boxplot) data for
 520 the weekly mean temperature using IBB with a 0.5°C temperature increase in Buan, South Korea.
 521 Boxes display the interquartile range (IQR), and whiskers extend to the extrema (i.e., maximum
 522 and minimum). The horizontal lines inside the boxes depict the median of the data. Note that the
 523 mean and maximum of the simulated data are increased significantly compared with the
 524 corresponding observed data, while the minimum of the simulated data is slightly increased and
 525 the standard deviation of the simulated data agrees with that of the observed data due to the
 526 variance inflation, as in Eq. (9).

527



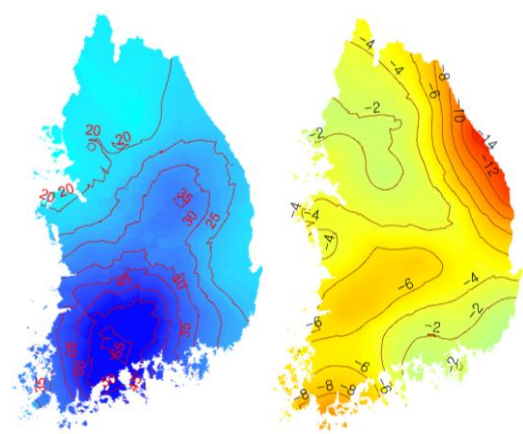
528
 529 Figure 9. The mean precipitation differences of the observed and simulated data (a) for the weekly
 530 precipitation in Buan with a 0.5°C mean temperature increase, (b) for the seasonal precipitation of
 531 all 54 stations with a 0.5°C mean temperature increase and (c) for a 1.0°C mean temperature
 532 increase. Note that indicates the mean of the simulated precipitation data for weekly (a) or seasonal
 533 (b and c).

534



(a) Winter

(b) Spring



(c) Summer

(d) Autumn

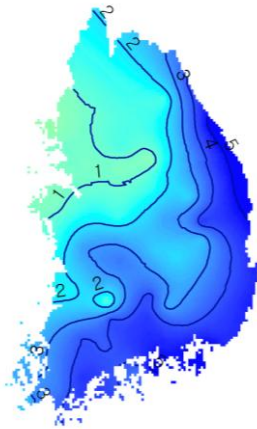
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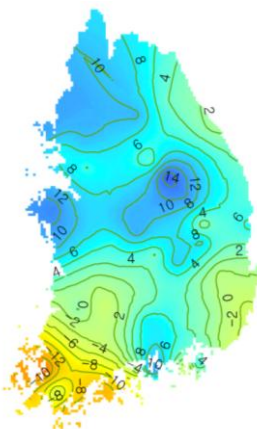
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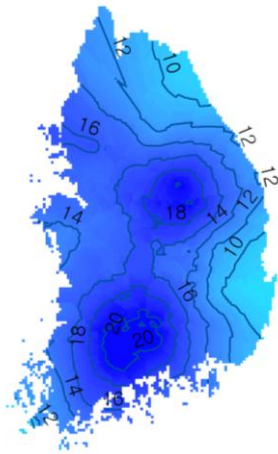
539 Figure 10. Spatial distributions in South Korea of the mean difference in seasonal precipitation
 540 (mm) with a 1.0°C increase in mean temperature. Note that the scale for the summer distribution
 541 is different from the other seasons, the 95% significance intervals are different at each station and
 542 the mean values of the significance intervals are ± 5.38 , ± 15.04 , ± 29.94 , and ± 4.84 for Winter
 543 (December, January, February), Spring (March, April, May), Summer (June, July, August), and
 544 Autumn (September, October, November), respectively.



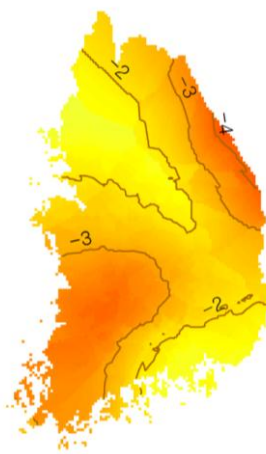
(a)Winter



(b)Spring



(c)Summer



(d)Autumn

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547

548

549 Figure 11. Spatial distribution of mean difference of seasonal precipitation (mm) with 0.5°C
 550 increasing mean temperature in South Korea. Note that the scale of summer is different from the
 551 other seasons and the 95% significance intervals are different at each station and the mean values
 552 of the significance intervals are ± 5.38 , ± 15.04 , ± 29.94 , and ± 4.84 for Winter (December, January,
 553 February), Spring (March, April, May), Summer (June, July, August), and Autumn (September,
 554 October, November) respectively.