Dear Dr. Elvidge (Referee, Geoscientific Model Development),

thank you for your reviewer report from 3 August 2016. We have accounted for the comments and suggestions in the revised manuscript version. Please find our replies to the particular comments in the following.

Sincerely,

Konrad Deetz and Bernhard Vogel

### Referee comments:

1) The VIIRS Night Fire (VNF) "flares only" dataset is not suitable for scientific applications. It is generated by stripping out VNF detections with either no temperature or temperatures under 1400K. This eliminates most biomass burning and ambiguous detections. The purpose of this is to provide a quick daily overview of global gas flaring activity. There are many times when a flare was detected in a single spectral band (usually M10 at 1.6 um), in which case the Planck curve cannot be fit and a temperature cannot be calculated. These detections have been lost in the dataset used by the authors. In addition, some flares are known to fluctuate in temperature and dip below 1400 K. These low temperature flaring events are also lost in the "flares only" daily summaries. The produce a more thorough analysis, the authors should work from the original daily VNF files. At best the "flares-only" version of the data provides a 'quick-and-dirty" depiction of global gas flaring.

For our work in the project Dynamics-aerosol-chemistry-cloud interactions in West Africa (DACCIWA) we wanted to have a consideration of gas flaring in our regional atmospheric model which includes the up-to-date characteristics of southern West Africa (SWA). The DACCIWA measurement campaign took place in June/July 2016 and for this time we need the flaring information for our model. Emission estimates for 2012/2013/2014 are not meaningful in our case, because the emissions are not constant from Also estimation year to year. your new (http://ngdc.noaa.gov/eog/viirs/download\_global\_flare.html) shows a decrease in flaring for Nigeria. To use older data would lead to overestimations.

The SWA emission inventory for flaring was not available when we started our research. The presented method is therefore our first approach to tackle the problem with the missing flaring emissions in our atmospheric chemistry simulations. Instead of using just constant emissions factors for flaring, we now have very regional information available.

We are concentrating on the description of the air pollution in our modelling system COSMO-ART and try to include all relevant emission sources. We are no experts in extracting the flaring sources from the general combustion sources detected by VIIRS Nightfire. Therefore we have relied on the "flares only" product published at <a href="http://ngdc.noaa.gov/eog/viirs/download viirs">http://ngdc.noaa.gov/eog/viirs/download viirs</a> flares only.html. Even if the data basis for our study is not perfect regarding VNF, there is a strong progress compared to the state before. We have changed our manuscript according to this problem. We have remarked, that the use of the "flares only" product is just a first approach and that this data contains greater uncertainties compared to the original VNF product. Future users of this parameterization can change the VNP input. The general method of the parameterization will not be affected by that.

2) The authors do not account for variations in cloud cover. This can be done based on the VIIRS Cloud product suite.

I see your point but our study focus is located to the creation of an emission dataset based on a VNF climatology rather than taking the VNF data day by day. In section 3.3.1 we describe the problem of flares that are masked by clouds (and the overall question whether the flare below the cloud is active or not) in detail and assess the uncertainty by using remote sensing cloud data from MSG and Aqua/AIRS. By deriving a flaring climatology (over two month), we are able to identify all flares (even if there are sometimes covered by clouds). With this climatological approach we get the mean emission strength of every flare (more precisely for every flare box). Therefore it is not necessary to account for the variations in cloud cover. Even if we would know, that a certain flare is masked by clouds at a certain day we don't know whether this flare is currently active and how large the radiant heat is. When we use our flaring climatology in our regional atmospheric model, all available flares are active at once with their mean emission strength.

3) The text should reference the following paper: Methods for Global Survey of Natural Gas Flaring from Visible Infrared Imaging Radiometer Suite Data (http://www.mdpi.com/1996-1073/9/1/14). We agree on that and have referenced the publication.

4) NOAA has global flaring data spanning 2012-2014 available at: http://ngdc.noaa.gov/eog/viirs/download\_global\_flare.html. There is a csv that contains locations and annual summaries of temperatures and radiant heat output of individual flares, normalized for cloud cover. The flared gas volume estimates are derived from an empirical calibration with CEDIGAZ reported flaring. It would be interesting to compare the NOAA results with those from the methods described in this paper.

From the xlsx file VIIRS\_Global\_flaring\_d.7\_slope\_0.331\_web.xlsx we have selected the 193 available Nigerian upstream flares and selected the flares which have a detection frequency greater than zero for 2014. We assume that "Avg. K" mean source temperature in K and "Ellipticity" means the radiant heat in MW. This data we have used as input for the parameterization presented in this study (with the same configuration). Finally we have integrated the volume stream of all Nigerian flare boxes from m-3 s-1 to m-2 y-1 and finally transformed it to bcm. The result is 8.55 bcm (271.0391 m-3 s-1). In the xlsx file the flared volume is estimated as 8.442995283 bcm for Nigeria in 2014. So if we use the same source temperature and radiant heat input as Elvidge et al. (2015) for Nigeria in 2014, we can reproduce the estimated flared volume with our method with a deviation below 1.3%.

Within our VNP data set for 2014 we estimate the flaring to 29.8 bcm. Regarding the uncertainty range of this estimation, the value is approx. by a factor of two higher than the other inventory. The uncertainty might result from the uncertainties in the estimation of the gauge pressure  $p_g$  and the fraction of the total reaction energy that is emitted as radiation f. In our flaring climatology we assume that all available flares are active at once with their mean emission strength, so this could lead to the higher values of the flared volume.

5) In the last sentence of the first paragraph, the text references the World Bank for a set of national flared gas volume estimates. The text should make it clear that these estimates were produced by NOAA using DMSP satellite data. There is a new set of estimates derived from VIIRS data at http://ngdc.noaa.gov/eog/viirs/download\_global\_flare.html.

We agree on that and have changed the manuscript accordingly. A remark to the availability of updated global flaring estimates for 2013 and 2014 at <a href="http://ngdc.noaa.gov/eog/viirs/download\_global\_flare.html">http://ngdc.noaa.gov/eog/viirs/download\_global\_flare.html</a> are mentioned in the manuscript.

Dear Dr. Elvidge (Referee, Geoscientific Model Development),

thank you for your reviewer report from 5 August 2016. We have accounted for the comments and suggestions in the revised manuscript version. Please find our replies to the particular comments in the following.

Sincerely,

Konrad Deetz and Bernhard Vogel

Referee comments:

1.) Being fully familiar with the flares only version of the VIIR Nightfire product I can certify that this product is not suitable for use in a scientific study. If the authors had contacted my team at the start of their study we could have explained this to them and directed them to the full VNF data files, which are suitable for use in scientific studies.

2.) NOAA does provide cloud state for each VNF detection - from the VIIRS cloud product suite. There are four states: confidently cloudy, probably cloudy, probably clear, confidently clear. But what is not recorded in the VNF files are the number of clear observations where the flare was not detected. The NOAA annual gas flaring data takes this into account. I dispute the authors contention that "it is not necessary to account for the variations in cloud cover."

3.) My overall impression is that these authors are willing to use data that are known to be flawed and ignore the effects of cloud cover variations in order to get a paper published without doing any addition work. If this journal is willing to publish papers with flaws like this disclosed - heaven help them.

The authors are puzzled about the different tenor in the two reviews we have achieved from you. We regret that you have the impression we are not willing to invest additional work for this study. Under point 4 of our reply from 3 August 2016 we followed your idea to compare your dataset with our study. Maybe one of your team can give us information on how to use the full VNF data files correctly and how to separate the flaring sources from other combustion sources (e.g. forest fires). With this data set, under consideration of the cloud correction, we will repeat our study.

M. Zhizhin provided us with the full VNF data for the relevant SWA countries in the time period of interest. The study has been repeated based on the new data and including the VIIRS cloud product.

Dear Mikhail Zhizhin (Referee, Geoscientific Model Development),

thank you for your reviewer report from 14 October 2016. We have accounted for the comments and suggestions in the revised manuscript version. Please find our replies to the particular comments in the following.

We have uploaded a revised manuscript which also includes the revision regarding the comments and suggestions of the first reviewer.

Sincerely Konrad Deetz and Bernhard Vogel

#### Referee comments:

In the paper a new method to model emissions from gas flaring is developed and validated on oil fields in Western Africa. The paper is a substantial contribution to the modeling science, and the approach is valid and motivating for further research. I have some comments on the presentation and details of the method which could be considered by the Authors before it is published.

0. I would recommend changing abbreviation VNP (VIIRS Nightfire Product) to commonly used VNF (simply VIIRS Nightfire) in the manuscript.

We agree on that and have changed the manuscript accordingly. For the general VIIRS Nightfire (including all combustion sources) we use the abbreviation "VNF" and for the extracted flaring information from VNF we use the abbreviation "VNF<sub>flare</sub>".

1. Formula (1) derives gas flow rate from flare radiative heat and temperature measured from satellite. It is a basis of the proposed model. However, it is taken from Appendix of regulating document by the German Environmental protection agency. This is technical, not scientific source. The derivation of the formula is not provided neither in the paper under review, nor in the cited document. The cited document has no source for the formula either. It is important to derive the formula (1) or to provide a scientific reference.

We agree on that. TA-Luft is a technical document and the equation is not well introduced there. Equation 1 of the manuscript originates from VDI 3782, 1985: Dispersion of Air Pollutants in the Atmosphere, Determination of Plume rise, Verein Deutscher Ingenieure, VDI-Richtlinien 3782 Part 3, Equation 24, https://www.vdi.de/richtlinie/vdi 3782 blatt 3-ausbreitung von luftverunreinigungen in der

atmosphaere\_berechnung\_der\_abgasfahnenueberhoehung/ (accessed: October 17, 2016). We have changed the citation accordingly. Although VDI 3782 (1985) is also a technical document, the derivation of the equation becomes clear. The heat flow M in MW is given by equation 1

$$M = c_p F (T_S - T_A), \tag{1}$$

where F is the flow rate in m<sup>3</sup> s<sup>-1</sup>,  $c_p$  the mean specific heat capacity of the emissions,  $T_S$  the source temperature and  $T_A$  the ambient temperature. VDI 3782 (1985) provides a value of the mean specific heat capacity of

$$c_p = 1.36 \cdot 10^{-3} MW s m^{-3} K^{-1}$$

which is derived for a pit coal firing but VDI 3782 (1985) denotes, that this can be used for other flue gases as well since potential deviations are negligible. (An explicit  $c_p$  value for gas flaring is not provided in the literature.) For the ambient temperature  $T_A$  we use 298.15K as a fixed value, representative for the tropical region. Within a sensitivity study regarding the influence of  $T_A$  on F we have used the mean heat flow and the mean source temperature of all flares in TP15 and varied the ambient temperature between 293K and 303K, as

a reasonable temperature range in the tropical regions. The resulting maximum difference in the heat flow is 0.0036 m3 s-1. Therefore we assume the errors using a fixed climatological value for the ambient temperature are negligible, but of course the user has to adapt the ambient temperature to the region he wants to apply the inventory. We have emphasized this in the manuscript. By using equation 1, the value for  $c_p$  and for  $T_A$ , the flow rate F in MW is given by:

$$F = M / (1.36 \cdot 10^{-3} (T_s - 298.15)).$$
<sup>(2)</sup>

2. Flare temperature used in the formula (1) is taken from instantaneous satellite measurement (VNF). It has a large variance depending on atmospheric conditions etc. I would recommend using mean flare temperature averaged over all cloud-free detections.

We see your point. This leads to a further source of uncertainty, because we cannot decide whether the spatial source temperature variations really results from the sources or from the atmospheric conditions. We think that this problem does not affect the climatological approach ( $E_{clim}$  in the revised manuscript) because for every detected flare the source temperature already is averaged over the two-month period of TP14 or TP15 before we calculate the emissions. We assume that this is a compromise between robustness and keeping the spatial variability of the flaring. To allow for consistency we now also use these temporal averages of source temperature and radiant heat for the daily resolved inventories ( $E_{obs}$  and  $E_{com}$  in the revised manuscript). Therefore all three inventories have the same underlying emission field and the difference is just related to number of flares that are active at a certain day. For  $E_{clim}$  all flares are active at once, for  $E_{obs}$  only the actual observed flares are active and for  $E_{com}$  the actual observed flares + the cloud covered flares (taken as active) are considered (com=combination). Nevertheless we have also included a further inventory in Tab. 5 that uses instantaneous input data to derive  $E_{clim}$  (first calculating the emissions for every single observation and then averaging the emissions temporally). This is given as " $E_{clim}$ , instant. input" and should allow further insight in the sensitivity/uncertainty.

3. The number 283 used in the formula (1) I believe stands for ambient air temperature at night? Is it a proper climatological value for Wester Africa?

Yes, the 283 refers to the ambient temperature. We agree that this value is not appropriate for the tropics. We have changed this value in the manuscript to 25°C (298.15K, also described in Comment 1 of this document). Owing to the change of the ambient temperature we have repeated our analysis to be consistent with this new value. The change from  $T_A = 283K$  to 298.15K lead to a slight increase in the emissions (e.g. for Fig. 9b the spatially integrated SWA emissions of TP14 increase from 651 to 658 t h-1).

4. Comments 1-3 may result in a wider variance of the proposed model output, and the model sensitivity analysis should be presented.

Regarding the ambient temperature (Comment 3) we have presented the maximum uncertainty in the heat flow as 0.0036 m3 s-1, which is also described in the manuscript. For the mean heat capacity of the emission  $c_p$  we do not have further information to assess the uncertainty. For considering the uncertainty in using temporal averages of source temperature and radiant heat instead of the instantaneous satellite observations, we have added a further emission inventory in Tab. 5 ("E<sub>clim</sub>, instant. input", for TP14 and TP15).

5. The Authors have made a considerable effort to take into account cloud conditions which can mask flare observations from space. Why not to use only cloud-free observation days, and to count detected/not detected flare cases to derive mean radiative heat?

By using the postprocessed flaring data (VNF<sub>flare</sub> in the revised manuscript which includes also a cloud mask) instead of the "Flaring only" product, it is straightforward to separate the flares into the categories (a) "cloud-covered", (b) "cloud-free and inactive" and (c) "cloud-free and active". By assuming that the cloud-covered flares are active with their mean emission strength, we can estimate the daily emissions via the sum of (a) and (c). To use only the cloud-free observation days would be problematic because SWA is a region with very extensive cloud cover (on average approx. 70% in the flaring area).

I would like to acknowledge that the Authors provide software sources and input data used in the study as the paper supplement. It is helpful for reproduction and reuse of their science and model.

List of relevant changes

1. Use flaring information from "VIIRS Nightfire Nighttime Detection and Characterization of Combustion Sources" instead of the quick-look data from "VIIRS Nightfire (Flares Only version). The study has been repeated with this data.

2. More detailed description and derivation of Equation 1.

3. Add missing references.

4. Based on the availability of the VIIRS cloud product, the strategy of assessing the uncertainty of flares masked by clouds has been changed. Three categories were defined: (1) cloud-free and active, (2) cloud-free and inactive, (3) cloud-covered and assumed to be active. This leads to different emission inventories. We have defined two inventories in addition to the climatology  $(E_{clim})$ : (a)  $E_{obs}$  which only consider the daily observations and (b)  $E_{com}$  which is a combination of Eobs and the emission from the cloud-covered flares.

5. Assessment of a further source of uncertainty regarding instantaneous VIIRS observations vs. averaged VIIRS observations. This assessment lead to a further emission inventory:  $E_{clim}$ , instantaneous input.

6. Additionally, the sensitivity of the flow rate calculation in Eq. 1 towards the ambient temperature has been assessed.

7. Correction of the ambient temperature in Eq. 1 to consider tropical conditions. For consistency the study has been repeated.

8. Update of Tab. 5 based on the revised analysis.

1 2

### Development of a new gas flaring emission data set for southern West Africa

- 3 Konrad Deetz, Bernhard Vogel
- 4

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### 8 **HIGHLIGHTS**

- 10 Development of a new gas flaring emission parameterization for air pollution modeling.
- 11 Combination of remote sensing observation and physical based combustion calculation.
- 12 Application to the significant gas flaring region southern West Africa.
- 13 Comprehensive assessing of the parameterization uncertainties.
- 14 Comparison with existing gas flaring emission inventories.
- 15

### 16 **Keywords:** 17

- 18 Gas flaring
- 19 Emission parameterization
- 20 Emission uncertainty
- 21 Pollution modeling
- 22 Carbon dioxide
- 23

#### 24 ABSTRACT 25

26 A new gas flaring emission parameterization has been developed which combines remote sensing 27 observations using VIIRS nighttime data with combustion equations. The parameterization has been 28 applied to southern West Africa, including the Niger Delta as a region which is highly exposed to gas 29 flaring. Two two-month datasets for June-July 2014 and 2015 were created. The parameterization 30 delivers emissions of CO, CO<sub>2</sub>, NO and NO<sub>2</sub>. A flaring climatology for both time periods has been 31 derived. The uncertainties owing to cloud cover, parameter selection, natural gas composition and 32 the interannual differences are assessed. Largest uncertainties in the emission estimation are linked 33 to the parameter selection. By using remote sensing cloud cover observations, a correction factor for the climatology was established to consider the effect of flares masked by clouds. It can be shown 34 35 that the flaring emissions in SWANigeria have significantly decreased by 3025% from 2014 to 2015. Existing emission inventories were used for validation. CO<sub>2</sub> emissions with the estimated uncertainty 36 in brackets of  $\frac{8 (\frac{12}{2})}{2.7 (\frac{3.6}{0.5})}$  Tg y<sup>-1</sup> for 2014 and  $\frac{5 (\frac{7}{12})}{2.0 (\frac{2.7}{0.4})}$  Tg y<sup>-1</sup> for 2015 are were derived. 37 The flaring Regarding the uncertainty range, the emission estimation within this study for June-July 38 39 2014estimate is in the same order of magnitude compared to existing emission inventories. For the same period in 2015 the emission estimation is one order of magnitude smaller in comparison to 40 41 existing inventories. with a tendency for underestimation. The deviations might be attributed to 42 uncertainties in a shortage in information about the derived flare gas flow rate combustion efficiency within southern West Africa, the decreasing trend in gas flaring or inconsistent emission sector 43 44 definitions. The parameterization source code is available as a package of R scripts. 45

- 46 **1. Introduction**
- 47

Gas flaring is a globally used method to dispose flammable, toxic or corrosive vapors to less reactive
 compounds at oil production sites and refineries. In regions of insufficient transportation
 infrastructure or missing consumers, flaring is also commonly applied.

51 CDIAC (2015a) estimated the global gas flaring emission of carbon dioxide to 267.7 million tons 52 (0.83% of total emissions) in 2008. Flaring and venting of gas significantly contributes to the 53 greenhouse gas emissions and therefore to the global climate change. The five countries with the 54 highest flaring amount in billion cubic <u>metermeters</u> (bcm) are Russia (35), Nigeria (15), Iran (10), Iraq 55 (10) and USA (5) (World Bank, 2012). These estimates were produced by National Oceanic and 56 Atmospheric Administration (NOAA) using Defense Meteorological Satellite Program (DMSP) remote

57 sensing data. Preliminary updates in global flaring estimates from NOAA for 2013 and 2014 are 58 available at http://ngdc.noaa.gov/eog/viirs/download\_global\_flare.html.

59 In recent time, especially with the development of remote sensing observation techniques (e.g. 60 Elvidge et al. (1997, 2013)), emissions from gas flaring moved in the focus of atmospheric research

61 involving the efforts in reducing the pollution and the waste of resources. The World Bank led the

62 initiatives "Global Gas Flaring Reduction Partnership" (GGFR) and "Zero Routine Flaring by 2030" to

63 promote the efficient use of flare gas.

64 Instead of relying on national statistics of gas production and consumption for estimating the flaring 65 amount, remote sensing techniques can estimate the flaring amount directly via multispectral data 66 (Elvidge et al., 2013). Elvidge et al. (2009) developed a 15 year dataset of global and national gas 67 flaring efficiency from 1994 to 2008 by using data from DMSP. Elvidge et al. (2015) presented 68 methods to derive global surveys of natural gas flaring using DMSP. For 2012 they have identified 69 7467 flares globally, with an estimated volume of flared gas of 143 ( $\pm$ 13.6) bcm. Doumbia et al. 70 (2014) combined DMSP with emission factors for flaring, to estimate the flaring emissions for SWA. 71 The satellite product Visible Infrared Imaging Radiometer Suite (VIIRS) Nightfire (Elvidge et al., 2013), 72 which is free available as "VIIRS Nightfire Prerun V2.1 Flares onlyNighttime Detection and 73 Characterization of Combustion Sources" (VIIRS, 2015) (VNP2015a) (VNF hereafter), is now the most 74 widely used product to derive flaring emissions from satellite imagery. By using VNPVNF, Zhang et al. 75 (2015) estimated the methane consumption and the release of  $CO_2$  from gas flaring for the northern

76 U.S. which agree with field data within an uncertainty range of  $\pm$  50%.

77 Also in the second largest flaring country Nigeria, the awareness of gas flaring increases. Nigeria 78 shows the fourth highest number of flare sites (approx. 300) worldwide after USA, Russia and Canada 79 (Elvidge et al., 2015). On gasflaretracker.ng the attention of the government, industry and society is 80 called to the flaring problem by interactive maps of flare infrastructure, amounts and costs. The 81 implications of gas flaring in Nigeria are far-reaching. It influences the environment by noise and 82 deterioration of the air quality (Osuji and Avwiri, 2005). Nwankwo and Ogagarue (2011) have 83 measured higher concentrations of heavy metals in surface water of a gas flared environment in Delta State Nigeria. Adverse ecolologicalecological and bacterial spectrum modifications by gas 84 85 flaring are indicated by Nwaugo et al. (2006). Gas flaring also causes acid rain which causes economic 86 burden via rapid corrosion of zinc roofs (Ekpoh and Obia, 2010) and causes retardation in crop 87 growth owing to high temperatures (Dung et al., 2008).

The project DACCIWA (Dynamics-aerosol-chemistry-cloud interactions in West Africa, Knippertz et al. (2015)) investigates the influence of anthropogenic and natural emissions on the atmospheric composition over SWA, including the flaring hotspot Nigeria, to <u>examinequantify</u> the <u>meteorological</u> and <u>socio-economic</u> effects <u>on meteorology and cloud characteristics</u>. To consider the SWA gas flaring emissions (e.g. in an atmospheric model), this study presents a method to derive emission fluxes by combining the state of the art flaring detection <u>VNPVNF</u> and the combustion equations of

- Ismail and Umukoro (2014) which does not use emission factors. The new parameterization is robust
   and easy to apply to new research questions according flexibility in the spatiotemporal resolution.
- The parameterization is presented in Section 2. Results of the application to SWA, including the spatial distribution of gas flaring, the emission estimation and the uncertainty assessment are investigated in Section 3. Section 4 places the emission estimates in the context of existing inventories. The results are summarized and discussed in Section 5.
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### 101 **2.** Parameterization of gas flaring emissions

102
 103 The new parameterization for gas flaring presented here, is based on <del>VNP (</del>VIIRS Nightfire <del>Prerun</del>
 104 V2.1 Flares onlyNighttime Detection and Characterization of Combustion Sources (VNF hereafter)
 105 and the combustion equations of Ismail and Umukoro (2014) (IU14 hereafter).

107 2.1 Remote sensing identification of gas flares

108 109 VIIRS (Visible Infrared Imaging Radiometer Suite) is a scanning radiometer for visible and infrared 110 light on board the sun-synchronous Suomi National Polar-orbiting Partnership weather satellite 111 (Suomi-NPP) (NASA, 2016). It can detect combustion sources at night (e.g. bush fires or gas flares) by 112 spectral band M10. To confirm these sources and to eliminate noise, the Day/Night Band (DNB), M7, 113 M8 and M12 are used in addition. By fitting these measured spectra to the Planck radiation curve, 114 background and source temperatures can be deduced. VNP is filtered to include only detections with 115 temperatures between 1600 K and 2000 K, which is believed to be an adequate estimation for average gas flares. Up to now no atmospheric correction is done (VIIRS, 2015 (VIIRS, 2015a). 116

- 117 The data is freely available as daily <u>cloud corrected</u> data from March 2014 to present. The files 118 include among others the location of the flares<u>combustion sources</u>, source temperature  $T_s$ , radiant 119 heat *H* and time of observation.
- 120 VNF does not distinguish between the different combustion sources (e.g. wild fires or flaring). To 121 extract the flaring information from VNF a postprocessing is necessary. For this study we have 122 decided for a two month period of observation. This allows a compilation of a flaring climatology in 123 terms of the locations and emissions and a robust estimation of uncertainty owing to cloud coverage and other parameters that have to be prescribed for IU14. We have selected the month June and July 124 125 because the gas flaring emission dataset will be used within the regional online-coupled chemistry model COSMO-ART (Vogel et al., 2009) during the measurement campaign of the project DACCIWA, 126 127 which takestook place in June/July 2016. This campaign includes airborne, ground based and remote 128 sensing observations of meteorological conditions and air pollution characteristics. COSMO-ART is 129 one of the forecasting models of the DACCIWA campaign and delivers spatiotemporal 130 aerosol/chemistry distributions. The data for June/July 2014 and June/July 2015 are used to allow 131 also for an interannual comparison, and to assess the uncertainty owing to changes in flare processes 132 (e.g. built-up or dismantling, increase or decrease in combustion). The dataset includes the countries 133 which can affect SWA with their flaring emissions, in particular lvory Coast, Ghana, Nigeria, Cameroon, Gabon, Congo, the Democratic Republic of the Congo and Angola. The extraction of the 134 135 flaring information from the VNF data (VNF<sub>flare</sub> hereafter) was realized by the Earth Observation Group of NOAA. Within VNF<sub>flare</sub> a csv file for every SWA flare is available, containing the flaring 136 137 history in June/July 2014 and 2015. For this study we use the location, source temperature and radiant heat. 138

For this study we use location, source temperature and radiant heat for days with sufficient satellite
 coverage over the research domain SWA with a focus on the Niger Delta.

141The authors want to point out that VNP delivers only a quick daily overview of global gas flaring142activity and is linked to uncertainties (e.g. flares in a single spectral band or with a source143temperature below 1400K cannot be detected). Instead of VNP, the original VIIRS files (available at144http://ngdc.noaa.gov/eog/viirs/download\_viirs\_fire.html)145parameterization presented in this study but this requires a preprocessing to separate the flaring

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#### 147

# 148 2.2 Emission estimation method149

sources from other combustion sources (e.g. wild fires).

The principle emission estimation methodology used in this study follows IU14. The gas flaring emissions are estimated based on combustion equations for incomplete combustion including six flaring conditions given in Tab. 1. The equations are introduced in detail in IU14 and are therefore not presented here. This section concentrates on the application of the method of IU14 to the VNPVNF<sub>flare</sub> data and the research domain-in SWA.

155

Tab.1. Reaction types for incomplete combustion of flared gas, depending on availability of sulfur in the flared gas and the
 temperature in the combustion zone which determines the formation of NO and NO<sub>2</sub>.

Reaction type	Sulfur in flared gas	Source temperature (K)	NO <sub>x</sub> formation
1	No	< 1200	no
2	Yes	< 1200	no
3	No	$1200 \le T_s \le 1600$	only NO
4	Yes	$1200 \le T_s \le 1600$	only NO
5	No	> 1600	NO and $NO_2$
6	Yes	> 1600	NO and $NO_2$

159

As input, IU14 needs the natural gas composition *C* of the fuel input of the flare, the source temperature  $T_s$  (temperature in the combustion zone), and the flare characteristics including combustion efficiency  $\eta$  (1 is complete combustion without Carbon monoxide formation) and availability of combustion air  $\delta$  (above 1 ismeans excess and below 1 ismeans deficiency). In addition we need the flow rate *F*, the gauge pressure of the fuel gas in the flare  $p_g$ , and the fraction of total reaction energy that is radiated *f*. The value for *f* is estimated by averaging a table of literature values for *f* given in Guigard et al. (2000). The IU14 input is summarized in Tab. 2.

167

168 **Tab.2.** Variables and parameters needed for IU14 or for deriving the fluxes of the air pollutants 169

Parameter	Description	Reference	Unit
С	Natural gas composition	Sonibare and Akeredolu (2004)	%
$T_s$	Source temperature	VNPVNF <sub>flare</sub> (VIIRS, <del>2015</del> 2015a)	К
η	Combustion efficiency	0.8 (IU14)	-
δ	Availability of combustion air	0.95 (IU14)	-
Н	Radiant heat	VNPVNF <sub>flare</sub> (VIIRS, 20152015a)	MW
F	Flow rate	VNPVNF <sub>flare</sub> (VIIRS, 2015), TA-Luft	m³ s⁻¹
		<del>(1986<u>2</u>015a), (VDI 3782, 1985</del> )	
$p_g$	Gauge pressure	34.475 (API, 2007)	kPa
$\vec{f}$	Fraction of radiated heat	0.27 (Guigard et al., 2000)	-

- 171 The natural gas composition is taken from Sonibare and Akeredolu (2004). They have measured the
- 172 molar composition of Nigerian natural gas in the Niger Delta area for ten gas flow stations. For this
- 173 study we have calculated the average over these stations and merged the data according their
- 174 number of carbon atoms (Tab. 3). H<sub>2</sub>S fraction is rather low because it was detected only in two out
- 175 of the ten flow stations.
- 176
- Tab.3. Molar composition of natural gas in Niger Delta (Nigeria) based on the measurements of Sonibare and Akeredolu
   (2004), averaged over ten flow station. The hydrocarbons are merged according to the number of C atoms.

Constituent	Fraction (%)	
Methan <u>Methane</u> (CH₄)	78.47	
EthanEthane (C₂H <sub>6</sub> )	6.16	
Propane (C <sub>3</sub> H <sub>8</sub> )	5.50	
<mark>Butan</mark> Butane (C₄H <sub>10</sub> )	5.19	
Pentane (C₅H <sub>12</sub> )	3.95	
Hexane $(C_6H_{14})$	0.36	
Carbon dioxide (CO <sub>2</sub> )	0.305	
Nitrogen (N <sub>2</sub> )	0.06	
Hydrogen sulfide (H <sub>2</sub> S)	0.005	

180

The source Temperature  $T_s$  is taken from  $\frac{VNPVNE_{flare}}{VNP_{flare}}$ . The combustion efficiency  $\eta$  was set to 0.8 and 181 the availability of combustion air  $\delta$  to 0.95. IU14 remarked, that the reaction condition for flaring of 182  $\eta \gg 0.5$  and  $\delta \ge 0.9$  should be the norm in regions, where the effective utilization of this gas is not 183 184 available or not economically. Strosher (2000) indicateindicates a combustion efficiency of solution gas at oil-field battery sites between 0.62 and 0.82, and 0.96 for flaring of natural gas in the open 185 186 atmosphere under turbulent conditions. EPA (1985) shows combustion efficiencies between 0.982 187 and 1 for measurements on a flare screening facility. Section 3.3.2 will shed light on the uncertainty 188 which arises from  $\eta$  and  $\delta$  via a parameter sensitivity study. The authors strongly recommend a 189 careful selection of  $\eta$  and  $\delta$  since unrealistic combinations (e.g. higher combustion efficiencies with 190 rather low availability of combustion air) can lead to negative NO and NO<sub>2</sub> emissions.

191 The flow rate, gauge pressure and fraction of radiated heat are not included in the parameterization 192 of IU14 but are necessary to derive the mass emission rates which can be used as emission data for 193 an atmospheric dispersion model.

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The flow rate F (m<sup>3</sup> s<sup>-1</sup>) is estimated by derived from Eq. 1 (TA-Luft, 1986 VDI 3782, 1985)

$$F = M/(1.36 \cdot 10^{-3} (T_{\rm s} - 283)), \tag{1}$$

$$F = M/(c_p (T_S - T_A)), \tag{1}$$

where *M* is the heat flow in MW-and-,  $c_p$  the mean specific heat capacity of the emissions,  $T_S$  the source temperature and  $T_A$  the ambient temperature. VDI 3782 (1985) provides a value of the mean specific heat capacity of

$$c_p = 1.36 \cdot 10^{-3} \, MW \, s \, m^{-3} \, K^{-1} \tag{2}$$

204which is derived for a pit coal firing but VDI 3782 (1985) denotes, that this can be used for other flue205gases as well since potential deviations are negligible. For the ambient temperature  $T_A$  we use206298.15K as a fixed value, representative for the tropical region. Within a sensitivity study regarding207the influence of  $T_A$  on the heat flow, we have used the averaged heat flow and source temperature

208of all flares within the time period June/July 2015 and varied the ambient temperature between209293K and 303K, as a reasonable temperature range in the tropical regions. The resulting maximum210difference in the heat flow is 0.0036 m3 s-1. Therefore we assume that the uncertainties using a fixed211climatological value for the ambient temperature are negligible. For the application of this inventory212to other regions the ambient temperature might be adapted. By using Eq. 1 and 2 the heat flow F can213be derived as

$$F = M / (1.36 \cdot 10^{-3} (T_s - 298.15)), \tag{3}$$

216 with  $T_s$  in K.

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217 We assume that the emitted heat flow M is equal to the total reaction energy of the flare. 218 VNPVNF<sub>flare</sub> only detects the energy fraction that is radiated H and not the total energy M. By using 219 the radiant heat H (observed by VNPVNF<sub>flare</sub>) and the factor f (fraction of H to the total reaction 220 energy, Guigard et al., 2000), we estimate M as  $H \cdot 1/f$ . For the source temperature  $T_S$  we use the 221 VNP observationVNF<sub>flare</sub> observations.

222 The estimation of the fuel gas density, which is necessary to transform the flow rate F into an 223 emission, is problematic due to the lack of data concerning the technical setup of the SWA flares. We 224 assume that the dominating flare type is a low-pressure single point flare. Bader et al. (2011) pointed out that these flares are the most common flare type for onshore facilities that operate at low 225 226 pressure (below 10 psi (69 kPa) above ambient pressure) and API (2007) remarks that most subsonic-227 flare seal drums operate in the range from 0 psi to 5 psi (34 kPa). Therefore we have decided for a gauge pressure  $p_g$  of 5 psi (34 kPa) above ambient pressure. Via Eq. 23 we can calculate the fuel gas 228 229 density  $\rho_f$ 

> $\rho_{f} = p_{f} / (R / (M_{f} T_{a})),$ (2)  $\rho_{f} = p_{f} / (R / (M_{f} T_{A})),$ (3)

where  $p_f$  is the fuel gas pressure as the sum of ambient pressure (10.1325 kPa, taken as const) and gauge pressure  $p_g$ . R is the universal gas constant,  $M_f$  the molar mass of the fuel gas and  $T_{\overline{\alpha}}T_A$  the ambient temperature (293298.15 K, taken as const). Finally, the emission E (kg s<sup>-1</sup>) of a species i is given by

$$E_i = \frac{m_i}{m_{total}} \rho_f F, \tag{34}$$

where  $m_i$  is the mass of the species *i* and  $m_{total}$  the total mass of the fuel gas, both delivered by the parameterization of IU14.

The combustion calculations within IU14 provide the species water, hydrogen, oxygen, nitrogen, 241 242 carbon dioxide, carbon monoxide, carbon dioxide, sulfur dioxide, nitrogen oxide and nitrogen dioxide. In the following only CO, SO<sub>2</sub>, NO and NO<sub>2</sub> the latter five are considered. However, no black 243 carbon or volatile organic compounds (VOCs) are considered by IU14, although they are not 244 negligible. Johnson et al. (2011) estimated the mean black carbon emission for a large-scale flare at a 245 gas plant in Uzbekistan to be 7400 g h<sup>-1</sup> and Strosher (1996) measured the concentration of 246 predominant VOCs 5 m above the gas flare in Alberta with 458.6 mg m<sup>-3</sup>. However, owing to the 247 248 missing representation of black carbon and VOCs in IU14, these compounds are not considered in 249 this study.

250 A flaring emission comparison between several days or averaging over a certain period is problematic 251 due to small variances in the VNP locations of the flares. This means even the same flare can be 252 detected on a slightly different position the next day, which makes an emission averaging for every 253 single flare difficult, especially in intensive flare areas. We bypass the problem by predefining a grid-254 and allocating the flares to this grid. By using the source code written in R (R Core Team, 2013) delivered by this study, the user can define the grid size independently. For calculating the average 255 256 over several days, the emissions for every single flare per day are calculated and summed up 257 according to their belonging to a certain grid box. This leads to one big point source per grid box. The 258 corresponding emissions are then averaged over the time period of interest for every grid box (flare 259 box herafter). Considering this approach within an atmospheric model, by selecting the same grid 260 configuration for the flaring emission data and the model, no loss of information occurs. 261 By using the source code written in R (R Core Team, 2013) delivered by this study, the user can define

262 the grid size independently (e.g. model grid) on which the flaring point sources are allocated.

### 263

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### 264 **3. Results** 265

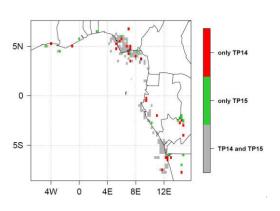
### 3.1 Spatial distribution of gas flaring in SWA

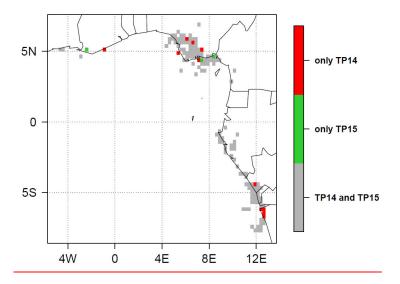
We have selected the two time periods June/July 2014 (TP14) and June/July 2015 (TP15) and omitted
all days without observations or with insufficient data coverage for VNP over SWA. This leads to 58
(48) observations for TP14 (TP15of VNF<sub>flare</sub> over SWA (61 observations respectively).

In the preparation of this work we have compared the estimated mean locations of the flares of TP14
with the Google Earth imagery (Google Earth, 2014) (not shown). Only the onshore flares are visible
in Google Earth. This visual verification reveals that 72% of the VNPVNE<sub>flare</sub> detected onshore flares
are visible in Google Earth. It is very likely that the hit rate is much higher since it is often the case
that the Google Earth image quality is not good enough for verification or the images are not up to
date. This comparison indicates that VNPVNE<sub>flare</sub> is a validan effective method to identify the flares in
SWA.

For the following analysis we have calculated the emissions for both time periods onallocated the
flares to a grid with a mesh size of 0.25° (28 km) from 108°S to 107°N and from 105°W to 1513°E- and
calculated the emissions for both time periods. A grid box with flaring is denoted as flare box
hereafter. Fig. 1 emphasizes the areas in which VNPVNE<sub>flare</sub> detects flares only in TP14 (TP15) in red
(green) color and in grey the areas with flaring in both periods.







**Fig.1.** Flaring area for TP14 and TP15. Red (green) boxes denote areas with flaring only for TP14 (TP15). For the grey areas, flaring is detected in both time periods.

Remarkable are the dominating flaring areas in the Niger Delta and the adjacent offshore regions in the Gulf of Guinea. Also in the coastal region of Gabon, Republic of the Congo, Angola and sporadically along the coast in Ghana and offshore of Ivory Coast, Ghana and Benin, flaring occurs. By comparing TP14 and TP15 more red than green areas are visible, especially in southern Nigeria, which indicates a reduction in the flaring area from 2014 to 2015. The red areas contribute 12% to the total CO<sub>2</sub> emissions of TP14. VNF<sub>flare</sub> detects 335 flares in 2014 and 312 flares in 2015 which means a reduction of about 7% (counted are those which deliver at least once a value for T<sub>s</sub> and H in the time period). 61% of that reduction is related to Nigeria. A decrease in CO<sub>2</sub> from 1994 to 2010, particularly in the onshore platforms is indicated by Doumbia et al. (2014).

300 Fig. 2 shows the density of flares (a) and the flaring activity (b) per flare box for TP15. The results are
 301 similar to TP14, therefore only the TP15 is displayed here.

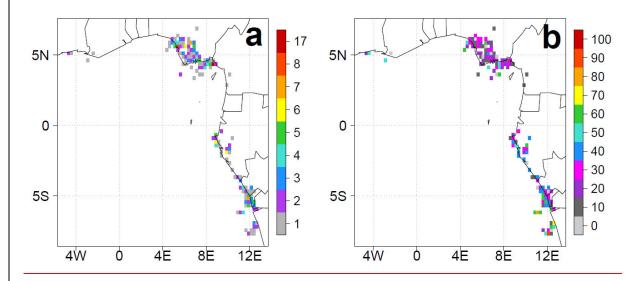


Fig.2. (a) Number of flares per flare box and (b) flaring activity (%) per flare box within TP15. A flaring activity of 100% means that every day on the 61 day period in June/July flaring was detected.

308 The highest flare density can be found offshore in the border area of Nigeria and Cameroon with 17 309 flares per flare box. The offshore flaring density is smaller than onshore (Fig. 2a) whereas the highest flaring activity can be found offshore (Fig. 2b). The mean active flare density, as the sum over all 310 detected flares in a box averaged over the time period, is shown in This could be linked to the 311 increased masking of flares by clouds over land. The large onshore flaring area of the Niger Delta 312 shows a comparable low flaring activity of 10-30%. Highest values can be found offshore of the 313 Democratic Republic of the Congo and Angola of 50-90%. How the interannual variability of flaring 314 315 reflects in the amount of flaring emissions is analyzed in section 3.3.4.

316 Fig. 2 for (a) TP14 and (b) TP15.



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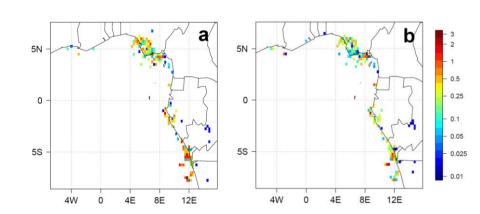
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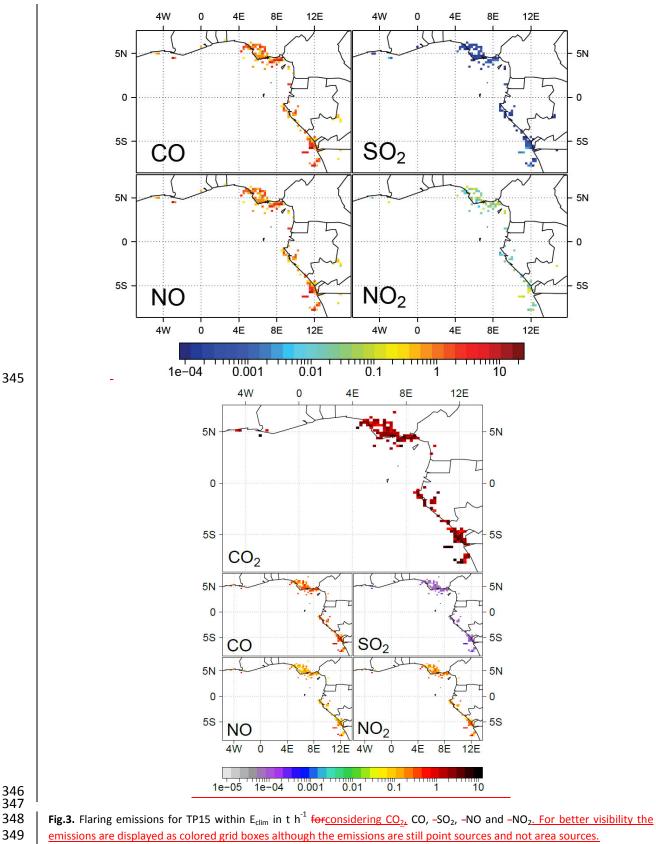


Eig 2-Mean active flare density (number of active flares per box) averaged over (a) TP14 and (b) TP15 in logarithmic scale

Fig. 2 shows to a a reduction in the active flare density in TP15 compared to TP14. 72% of the flaring area which TP14 and TP15 have in common, shows a reduction in TP15 about 48% on average. 28% of the common flaring area shows an increase in TP15 about 124%. Therefore it seems that the flaring intensity decreases in TP15 over large areas but simultaneously some flaring hotspots occurred, which are distributed along the SWA coast (not shown). Fig. 2, together with the variation of flaring emissions from TP14 to TP15 in Section 3.3.3, indicates the high year to year variations. This makes the use of past averaged conditions questionable, especially when certain episodes are studied.

#### 330 3.2 Emission estimation

For the emission estimation we have used a climatological approach (E<sub>clim</sub>). For every day with valid 332 333 data in TP14 and TP15 the emissions for all detected flares are calculated separately and allocated to 334 the predefined grid. The emissions are summed up in every flare box to have one joined flare per grid box. Finally the temporal average for every grid box is calculated averages of source temperature and 335 336 radiant heat over TP14 and TP15 respectively. were used to calculate the emissions. Therefore in this 337 approach all flares, detected in the time period, are active at once with their mean emission strength. This method has the advantage that most likely all flares in the domain are captured even if 338 339 a fraction of them is covered by clouds at certain days. However, this could lead to an emission 340 overestimation because not all available flares are active at once. This problem of separating 341 between flares which are not active and flares which are active but covered by clouds and therefore 342 not visible for VNPVNF<sub>flare</sub> is picked up again in Section 3.3.1. Fig. 3 shows the emissions of CO<sub>2</sub>, CO,  $SO_2$ , NO and  $NO_2$  in t h<sup>-1</sup> for TP15. 343



emissions are displayed as colored grid boxes although the emissions are still point sources and not area sources.

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351 Highest emissions are calculatedderived for carbon monoxidedioxide, followed by carbon monoxide, nitrogen oxidedioxide and nitrogen dioxideoxide. Sulfur dioxide shows lowest emissions since these 352 353 emissions do not depend on combustion processes but only on the natural gas composition (see Tab. 354 3) and the amount of flared gas (IU14). Due to the use of the averaged measurements of Sonibare and Akeredolu (2004), local variations of hydrogen sulfide concentrations in the natural gas cannot be taken into account. Hydrogen sulfide is the only source of sulfur in the flared gas and therefore determines the emission of sulfur dioxide. To assess this uncertainty, a sensitivity study with different hydrogen sulfide concentrations is given in Section 3.3.5.

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# 360 3.3 Estimation of uncertainties 361

362 In the following section the most relevant uncertainties are presented, together with approaches for 363 their assessment. This includes the uncertainty concerning the flare detection in the presence of 364 cloud cover, the uncertainty in the determination of the emitted heat flow H via the fraction of 365 radiated heat f, the uncertainty in the choice of the IU14 parameters and the changes in flare 366 operation from one year to another as well as the influence of the spatial variability of hydrogen 367 sulfide in the natural gas on the sulfur dioxide emissions. Apart from Section 3.3.4 all uncertainty 368 estimations are confined to TP15.

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# 370 3.3.1 Uncertainty due to cloud cover371

In this section we want to estimate the emission error due to cloud-covered flares and present a 372 373 method to derive daily emissions by considering the contribution of these masked flares. In Section 374 3.2 a climatological data set of flaring emissions (E<sub>clim</sub>) was derived. When using this data set we, in which all available flares are losing the active with their mean emission strength. This dataset 375 376 therefore does not include a day to day variation of the flaring emissions that. If an emission dataset 377 with a daily variability is delivered by VNP. Although daily satellite observations are 378 available required, the problem arises that usually parts of the scene observed by the satellite are 379 covered by clouds. In and therefore the following we will describe a method of how to derive daily 380 flaring emissions based on are likely underestimated. VNF flare includes the climatological emissions 381 (E<sub>clim</sub>), a threshold of cloud coverage (N<sub>th</sub>), and locations of all flares independent whether there are 382 active or not. This entity is illustrated by the actual detected flares closed dark grey pie in Fig. 4A and 383 4B. By comparing the flares which are observed/active at a certain day. This is illustrated 384 schematically by Fig. 4. and the total

385 The-closed grey pie in the lower layer of Fig. 4A gives the climatological number of flaring boxes in 386 flares, a separation between observed (green pie in Fig. 4A) and not observed (light grey pie in Fig. 4A) is possible. In addition VNF<sub>flare</sub> delivers a cloud mask for all of the research domain. At a certain 387 388 day only within-flare detections. Therefore it is possible to separate the green flaring boxes active flares are detected by VNP. The flaring boxes that are indicated in grey are those at which no active 389 390 flares were detected by VNP, either because they are light grey pie of the not observed flares in (a) 391 cloud-free and inactive or obscured by clouds. We now further separate this grey area by introducing 392 an empirical threshold value N<sub>th</sub> of cloud cover. In areas that belong to the grey fraction in Fig 4A, 393 where the cloud cover is above N<sub>th</sub>, we assume that the flares boxes are active and emit with their 394 climatological emission values (since there are no current observations available). Those flare boxes 395 are indicated by the dark blue color in Fig 4B. The light blue area indicates flare boxes where the 396 cloud cover is below Nth and where no flares are detected by VNP. For this area we postulate that all 397 flare boxes are inactive and consequently have zero emissions. Finally we calculate the total 398 emissions at a certain day for N<sub>th</sub>=50% (E<sub>50</sub>), 75% (E<sub>75</sub>) and 90% (E<sub>90</sub>) as the sum of the climatological 399 emissions in the dark blue area and the directly detected flares in the green area. (light blue pie in 400 Fig. 4B) and (b) cloud-covered and unknown flaring status (blue pie in Fig. 4B).

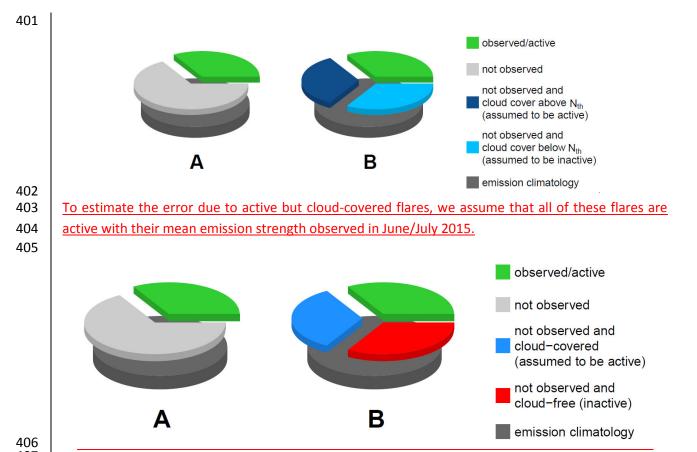


Fig.4. Pie charts illustrating the flaring emission uncertainty assessment due to cloud cover for TP15. The entiretyentity of

the flare boxesflares within the emission climatology (Eclim) is given as closed grey pie in the bottom of A and B. A

distinguishes between flare boxes in which flares which are detected/active at a certain day (green) and the complement of

undetected flare boxes flares (light grey). In B the light grey slice of A is separated in a cloud-covered (above cloud cover

<del>threshold N<sub>th</sub>, dark b</del>lue) and cloud-free (<del>below N<sub>th</sub>, light blue)<u>red) part</u> by using <del>remote sensing observations. Flare</del></del>

boxesthe cloud mask of VNF<sub>flare</sub>. Flares which are not detected by VNP<u>VNF<sub>flare</sub></u> and simultaneously show a cloud cover above

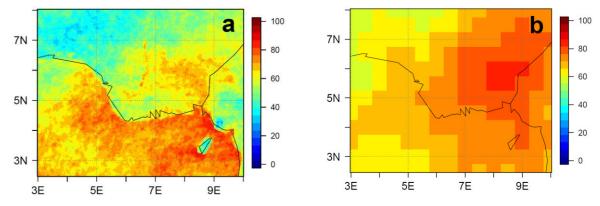
Nthy covered by clouds are taken as active. Flare boxesFlares which are not detected by WNPVNF flare and simultaneously show

a cloud cover below N<sub>th</sub>, are not covered by clouds are taken as inactive. For N<sub>th</sub> the values 50%, 75% and 90% are used. The

higher N<sub>th</sub> the smaller the dark blue slice in **B**.

Fig. To separate the light grey slice in Fig. 4A in covered and uncovered flare boxes, we used 5 418 419 illustrates the mean cloud cover exemplarily for the greater Niger Delta area using (a) instantaneous 420 cloud fractional cover (CFC) from the geostationary Meteosat Second Generation 3 (MSG3) (CM SAF, 421 2015, copyright (2015) EUMETSAT) for every day of TP15 around the time of VNPVNF observation 422 (Suomi-NPP overflight approx. at 1 UTC). This method is applied to all days of TP15 for every flare 423 box.) and (b) the sun-synchronous Aqua/AIRS (Mirador, 2016). To ensure a consistent timing between cloud observation and VNP observation, the spatial domain 424 425 was reduced with a focus on the Niger Delta area (see Fig. 5a) and the flares were allocated

- 426 according to the cloud data grid with a mesh size of 0.03°.
- 427

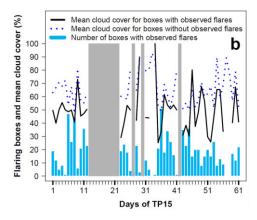


429 Fig.5. Fractional cloud cover (%) observed from (a) the geostationary MSG3 and (b) the sun-synchronous Aqua/AIRS,
 430 averaged over TP15 around the time of <u>VNPVNF</u> observation (approx. 1 UTC).

Fig. 5a shows that the onshore flaring area for TP15 is in mean covered with clouds by 50-70%. For the offshore flaring area it is even higher with 70-90%. Therefore it is very likely that flares are frequently masked by clouds and therefore not detected by <u>VNPVNF</u>. However, we suspect that the MSG3 cloud product underestimates (overestimates) the onshore (offshore) cloud cover when comparing with the findings of van der Linden et al. (2015). The high offshore coverage and the distinct land-water separation might be caused by overestimating low clouds in the presence of a warm and moist tropical ocean.

Fig. 5b shows a cloud climatology using Aqua/AIRS Nighttime data (Mirador, 2016). The Aqua/AIRS 439 440 climatology shows higher cloud cover over land and no distinct separation between water and land 441 surface. Both products identify the highest onshore cloud cover in the northeast of Port Harcourt 442 (4.8°N, 7.0°E) and have similar values in the Nigerian offshore region (containing the offshore flares) 443 of about 70-80%. The major difference in the climatologies appears onshore between 4.5°N and 6°N. This area includes the majority of the Nigerian onshore flares. Although it is not the aim of this study 444 to identify the most reliable cloud climatology for SWA, it has to be considered that MSG3 likely 445 underestimates the mean cloud cover over the Nigerian onshore flares up to 30%. This reveals a 446 447 relatively high uncertainty in the estimation of nocturnal low cloud coverage from remote sensing. However, in the following the cloud climatology derived from MSG3 (Fig. 5a) is used since Aqua/AIRS 448

449 cannot provide the full spatial coverage for every day (due to the sun synchronous orbit of
 450 Aqua/AIRS).



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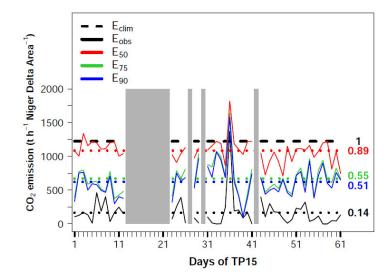
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Fig.6-Number of boxes with detected flares per day (blue bars) and the mean fractional cloud cover for the boxes with (without) detected flares as black solid (blue dotted) line (using MSG3, compare Fig. 5a). For the calculation of the latter, the cloud cover of the non-active flare boxes within-E<sub>clim</sub> are averaged (compare Niger Delta area in Fig. 2b). The grey shaded areas are omitted due to lack of VNP observation.

- 459 Fig. 6 shows the number of grid boxes with active flares per day in TP15-as, separated in the categories: cloud-free/active (green), cloud-free/inactive (red) and cloud-covered (blue-bars. The 460 grey areas indicate). Flares with no or incomplete data gapsare coded in VNPblack. E<sub>clim</sub> includes 185 461 462 flare boxes according to the domain in Fig 5a. For TP15 not more than 51 flare boxes are detected at 463 once. In 312 flares which are at least once active in TP15. On average only 826% of the total flaring area is active at once. As expected the temporal evolution of the flare boxes and the, 9% is verifiable 464 465 inactive and 63% is cloud-cover for these boxes (black solid line in Fig. 6) shows an anticorrelation. The highest number of flare boxes at day 36 is reached in a period of a comparatively low cloud 466 cover. The mean cloud cover for the non-active flare boxes of E<sub>stim</sub> (blue dotted line in Fig. 6), is in 467 general higher than for the active flare boxes which implies that the cloud cover reduces the VNP 468 detections. Fig. 6 also reveals that it is not suitable to use the strict cloud free condition for the 469 470 separation in Fig. 4B because nearly all of the boxes would be assigned to the dark blue cloud 471 covered fraction and the resulting emissions would be nearly the same as E<sub>clim</sub>.
- However, it has to be considered that the light points of flares are extremely small-scale signals 472 473 (1/5000 of the VNP pixel, Zhang et al. (2015)) and even for an almost completely closed cloud deck 474 VNP detections are possible.
- 475 The climatology E<sub>clim</sub> is the reference for this study. In addition we define E<sub>abs</sub> which only considers the actually observed flares per day. E<sub>so</sub> is defined as the combination of actually observed flares and 476 477 cloud-covered-flares (see Fig. 4) with a. By taking into account only the cloud-cover threshold-free 478 information instead of 50%. E25 (Enn) is equal to E50 but uses a cloud cover threshold the climatological 479 approach of 75% (90%).
- 480 To emphasize the difference between the different emission estimates, Fig. 7 shows the daily 481 emissions of CO<sub>2</sub> for TP15 as a spatial sum over the Niger Delta area (see Fig. 4a). In contrast to E<sub>rlim</sub> (black dashed line), E<sub>obs</sub>, E<sub>50</sub>, E<sub>25</sub> and E<sub>90</sub> (solid lines) have a temporal variation within TP15. 482
- 483

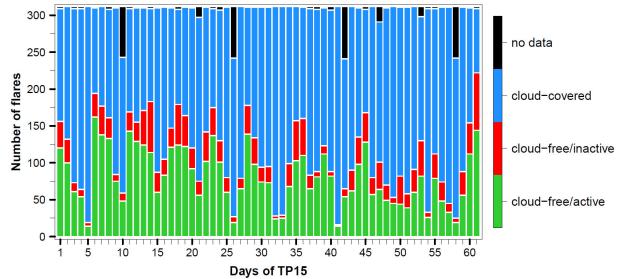
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Fig.7. Daily CO<sub>2</sub> emissions (kg h<sup>-1</sup>) within TP15 from flaring summed up over the Niger Delta area defined in Fig. 4a for the estimates: Entry (climatology, black dashed line), Eobs (VNP observations, black solid line), Est (combination VNP observations and the climatology for a cloud cover threshold of 50%, red solid line), E25 (as E50 but for a cloud cover threshold of 75%, green solid line). Example for a cloud cover threshold of 90%, blue solid line). The dotted lines denote the spatiotemporal, On average of East, Eso, Eso, Eso, The numbers on the right hand side show the ratios of the spatiotemporal averages East, Eso, Eso, Eso, Eso, eso, towards Eclim. The grey shaded areas are omitted due to lack of VNP 492 observation.

494  $E_{clina}$  delivers a daily CO<sub>2</sub> emission of about 1250 t h<sup>-1</sup> within the Niger Delta area. The pure daily VNP observations within Easter (black solid line) show only 14% of Easter emissions (numbers on left hand side 495 of Fig. 7) on average (black dotted line). The emissions from VNP observations together with the 496 climatology for the cloud threshold of 50% within E<sub>so</sub> (red solid line) is closest to the climatology (89% 497 498 of Eclim, red dotted line). The high overall cloud cover within the domain (compare with blue dotted 499 line in Fig. 6) together with the relative low cloud cover threshold leads to the result, that nearly the 500 complete climatology is used for E<sub>50</sub> and therefore the difference to E<sub>clim</sub> is small. The emissions from VNP observations together with the climatology for the cloud threshold of 75% and 90% within Egg 501 and E<sub>no</sub> (green and blue solid line) shows only small deviations but are significantly reduced in 502 comparison to Eclim (55% and 51% of Eclim, green and blue dotted line). Day 36 of TP 15 shows highest 503 504 emissions in E<sub>obs</sub>, E<sub>50</sub>, E<sub>75</sub> and E<sub>90</sub>, owing to the combination of low cloud cover and high flaring 505 activity (compare with Fig. 6). Regarding the uncertainty in the cloud cover climatology (compare Fig. 506 5a and Fig. 5b), the emissions of E<sub>50</sub>, E<sub>75</sub> and E<sub>90</sub> might be underestimated. The 63% of the flares are 507 not considered at a certain day. By assuming that all of these cloud-covered flares are active, a 508 remarkable underestimation of the cloud cover in the onshore flaring area could lead to an 509 unjustified increase in flare boxes below N<sub>th</sub> and therefore to a reduced number of active flares per 510 daycan be expected.



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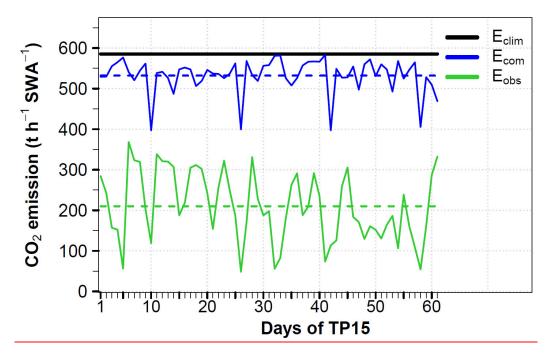
**Fig.6.** Number of flares per day in TP15 which are cloud-free and active (green), cloud-free and inactive (red) and cloudcovered (blue). Flares with no or incomplete data are denoted in black. The color coding follows Fig. 4B. Considered are the 312 flares which deliver at least once a value for  $T_s$  and H in TP15.

In addition to E<sub>clim</sub> two further emission inventories are introduced: E<sub>obs</sub> only considers the actual daily observed flares (linked to the green flares in Fig. 6). To consider also the contribution of active but cloud-covered flares, E<sub>com</sub> combines the green and the blue flares of Fig. 6.

<u>To allow for consistency, all three inventories use the emissions derived from the flare specific temporal averages of the source temperature and the radiant heat over TP14 and TP15 respectively.</u>

We avoid calculating the emissions from instantaneous source temperatures because this is linked to
 high uncertainty depending on the atmospheric conditions (Mikhail Zhizhin, personal
 communication). The temporal averages allow for robustness. Therefore the three inventories only
 differ in the selection of the active flares per day but not in the underlying emissions. E<sub>clim</sub> uses all
 flares at a certain day, E<sub>obs</sub> considers only the flares which are cloud-free and active and E<sub>com</sub>

527 considers E<sub>obs</sub> plus the cloud-covered flares, by assuming that all of the cloud-covered flares are
 528 active. Nevertheless we have included a further inventory in Tab. 5 which uses instantaneous source
 529 temperature and radiant for the emission derivation (E<sub>clim</sub>, instant. input) to assess the differences
 530 towards the averaged input. Fig. 7 shows the total CO<sub>2</sub> emissions of the SWA area from E<sub>clim</sub> in black,
 531 from E<sub>obs</sub> in green and from E<sub>com</sub> in blue.



**Fig.7.** Daily CO<sub>2</sub> emission estimations (t h<sup>-1</sup>) within TP15 from flaring, summed up over the SWA area as denoted in Fig. 1 for the three emission inventories:  $E_{clim}$  (climatology, black solid line),  $E_{obs}$  (daily VNF<sub>flare</sub> observations, green solid line and temporal average as green dashed line) and  $E_{com}$  (sum of daily VNF<sub>flare</sub> observations and emissions from cloud-covered flares, blue solid line and temporal average as blue dashed line). The periodical drop of the blue line is linked to reduced data coverage (compare with black bars in Fig. 6).

The dashed lines denote the temporal averages of  $E_{obs}$  and  $E_{com}$ . On average  $E_{com}$  is only 9% smallerthan  $E_{clim}$  which is assumed to be in the range of uncertainty. Therefore both inventories areequitable in this study. The user can decide whether a temporal resolved or a climatologicalapproach fits best to their research question.

**i**45The emissions of  $E_{obs}$  are strongly reduced (64%) compared to  $E_{clim}$  as expected. The use of  $E_{obs}$  would**i**46significantly underestimate the emissions and is therefore not appropriate for an application. Since**i**47 $E_{obs}$  does not take into account cloud-covered flares at all and  $E_{com}$  in contrast sees all cloud-covered**i**48flares as active, the difference between these inventories can be used to assess the uncertainty**i**49arising from flares masked by clouds. Fig. 7 shows a mean difference between  $E_{obs}$  and  $E_{com}$  of about**i**5061%. Therefore while using  $E_{obs}$  as a flaring emission inventory in an application, an underestimation**i**51of the emissions of 61% has to be considered.

These emission estimations contain different information.  $E_{clim}$  includes all flares of the domain despite cloud coverinvariant but can overestimate the emissions.  $E_{obs}$  shows the <u>VNPVNF<sub>flare</sub></u> reality, including a temporal development, but cannot consider the cloud-covered flares.  $E_{507}$   $E_{75}$  and  $E_{90}$ combine<u>E<sub>com</sub></u> combines the flare location<u>climatological</u> information of  $E_{clim}$  for flares which are not observable at a certain time and the full temporal resolution of <u>VNPVNF<sub>flare</sub></u> in  $E_{obs}$  by using cloud observations. However this approach is based on the assumption that all cloud-covered flare boxes are active, which is also linked to high uncertainty. Additionally  $E_{507}$   $E_{75}$  and  $E_{90}$  depend on the availability of a longer VNP observational dataset. The ratios of the spatiotemporal means of E<sub>obs</sub>, E<sub>50</sub>,
 E<sub>75</sub> and E<sub>90</sub> to the spatial mean of E<sub>clim</sub> (as denoted by the numbers in Fig. 7) are used as correction
 factors (CF) for E<sub>clim</sub> in the following (see Tab. 4). E<sub>clim</sub> is taken as the reference (CF=1).

### 561 562 563 564 565

566

**Tab.4.** Emission estimations including information about flaring (daily observation and climatology) and cloud cover observation. The correction factors (CF) are derived for TP15 from a spatiotemporal emission mean in the Niger Delta area (2.5°N-8°N, 3°E-10°E) and refer to E<sub>clim</sub>.

Name	Emission estimate	CF for E <sub>clim</sub>
E <sub>clim</sub>	Climatology (reference)	1
<del>E<sub>obs</sub></del>	Observed flares	<del>0.14</del>
<del>E</del> 50	Observed flares + climatology (N <sub>th</sub> = 50%)	<del>0.89</del>
<del>E</del> 75	<del>Observed flares + climatology (N<sub>th</sub> = 75%)</del>	<del>0.55</del>
<mark>€</mark> 90	Observed flares + climatology (N <sub>th</sub> = 90%)	<del>0.51</del>

567

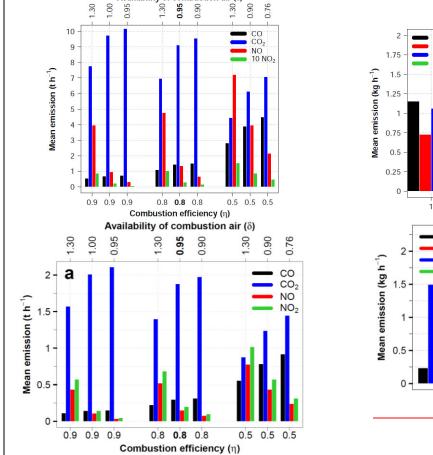
These CF are a simple method to include the information of E<sub>obs</sub>, E<sub>50</sub>, E<sub>75</sub> and E<sub>90</sub> into E<sub>clim</sub> by 568 569 multiplying E<sub>clim</sub> with the corresponding correction factor. In this case the same 185 flare boxes of 570 Eclim are used but with an emission strength reduced to the averaged conditions of Eobs, E50, E75 and Ener. This approach is based on the assumption that the correction factor, deduced for the Niger Delta 571 572 area, is valid for the whole domain specified in Section 3.1. This assumption seems to covered flares 573 are active, which can be justified since the Niger Delta area containsseen as an estimation upwards. 574 Therefore the most of the gas flares in the domainlikely amount of emissions is expected between 575 Eobs and Ecom.

#### 576

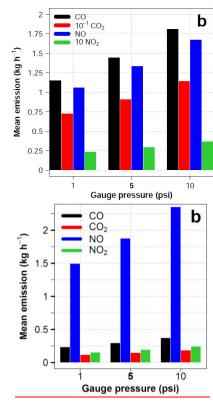
# 577 *3.3.2 Uncertainty due to IU14 input parameters* 578

579 To assess the uncertainty which arises from the combustion efficiency  $\eta$  and the availability of combustion air  $\delta$ , a sensitivity study has been carried out. The exact values for the SWA flares are 580 581 unknown and very likely highly variable from one flare to another, depending on the flare type and operation. Fig. 8a shows the flare flaring emissions averaged over SWA and TP15 for CO, CO<sub>2</sub>, NO and 582 NO<sub>2</sub>. The parameters  $\eta$  and  $\delta$  are varied referring to IU14. A complete combustion ( $\eta = 1$ ) does not 583 produce CO emissions since all carbon is transformed to CO<sub>2</sub> (not shown). With decreasing  $\eta$  and  $\delta$ , 584 585 the CO and CO<sub>2</sub> emissions increase. Concerning CO we assume the lower limit for  $\eta = 0.9$  and 586  $\delta = 1.3$  (left of Fig. 8a) and the upper limit for  $\eta = 0.5$  and  $\delta = 0.76$  (right of Fig. 8a). The values 587 used for this study are located in the center of Fig. 8a.8a (printed in bold). By taking the latter as 588 reference, the lower (upper) limit leads to a decrease (increase) in CO emission of -63% (+210208%). 589 For CO<sub>2</sub> we derived an lower (upper (lower) limit of +38% (-72-53% (+12%).

590 A higher combustion efficiency or a higher availability of combustion air allows an enhanced 591 formation of NO and NO<sub>2</sub>. Therefore NONO<sub>x</sub> emissions increase (decrease) with decreasing  $\eta$ -. In 592 contrast these emissions decrease with an increase in the combustion efficiency ( $\delta$ ). The higher the 593 efficiency the more oxygen is forming CO<sub>2</sub> instead of NO<sub>x</sub>. We assume the lower limit for  $\eta = 0.9$  and 594  $\delta = 0.95$  and the upper limit for  $\eta = 0.5$  and  $\delta = 1.30$ . Taking again the central parameter set of Fig. 595 8a as reference, the lower (upper) limit leads to a decrease (increase) in NO emission of -77% 596 (+44176% (+420%).



Availability of combustion air (8)



**Fig.8.** Flaring emissions (kgt h<sup>-1</sup>) spatiotemporally averaged over SWA and TP15 depending on (a) combustion efficiency  $\eta$  and availability of combustion air  $\delta$  for a gauge pressure of 5 psi and (b) gauge pressure (psi) for the setup of  $\eta = 0.8$  and  $\delta$  which is used for this study (emphasized in bold).= 0.95. SO<sub>2</sub> is not shown because it does not depend on  $\eta$  or  $\delta$ .

The emissions of NO<sub>2</sub> are comparatively low owing to the source temperature which is in general
 lower than the NO<sub>2</sub> formation threshold of 1600 K.

605 For NO<sub>2</sub> the emission decrease (increase) is -76% (+417%).

606 In addition, Fig. 8b shows the emissions depending on the gauge pressure for 1 (lower limit), 5 and 607 10 psi (upper limit) (7, 34 and 69 kPa respectively) for  $\eta = 0.8$  and  $\delta = 0.95$ . RegardingUsing 5 psi as 608 the reference, the lower (upper) limit leads to a decrease (increase) in CO emissions of -21% (+2620% 609 (+25%).

610Fig. 8 emphasizes that the technical conditions of flaring crucially influence the emission strength and611that the emissions are more sensitive towards  $\eta$  and  $\delta$  than towards the gauge pressure.

612

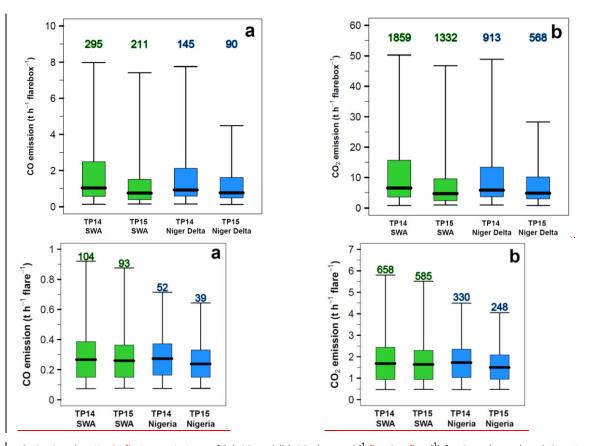
#### 613 *3.3.3 Uncertainty due to the fraction of radiated heat* 614

To estimate the uncertainty in the fraction of radiated heat f (see Tab. 2), we have used the standard deviation of the literature values given in the appendix of Guigard et al. (2000) in addition to the mean value of f = 0.27. This leads to a domain of uncertainty for the value f of  $(^{0.38}/_{0.16})$ . Therefore the <u>VNPVNF<sub>flare</sub></u> observed radiant heat is multiplied with the factor 1/f of 3.7 ( $^{6.2}/_{2.6}$ ).

- 620 3.3.4 Interannual variability
- 621

The differences in flaring between TP14 and TP15, indicated in Fig. 1 and Fig. 2, are quantified in this section according to the emissions of CO (Fig. 9a) and  $CO_2$  (Fig. 9b). The boxplots include all flaring boxesflares for the two domains SWA (green) and the Niger Delta areaNigeria (blue). The numbers above indicate the integrated emissions per hour and area in tons.





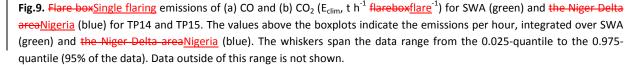
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The emissions of CO<sub>2</sub> are 6.3 times higher than the CO emissions. For SWANigeria (blue boxplots) the
mean value of emissions is statistically significant lower for TP15 compared to TP14 (Wilcoxon-MannWhitney rank sum test with a significance level of 0.05). For the Niger Delta area SWA the emission
meansaverages show no significant difference. The significant different mean values for SWA
emissionsNigeria emphasize the relevance of using a flaring dataset which is up to date to reduce
uncertainties arising from deviations in flare locations or flaring processes.

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# $\begin{array}{ll} 640 \\ 641 \end{array} \qquad \textbf{3.3.5 Uncertainty due to spatial variability in $H_2$} \\ 641 \end{array}$

Since hydrogen sulfide (H<sub>2</sub>S) is the only sulfur source in the flared gas, it determines the emission of sulfur dioxide. The natural gas composition measurements from the ten flow stations given in Sonibare and Akeredolu (2004) contain only two stations with nonzero H<sub>2</sub>S content. Therefore averaging over the ten stations (see Tab. 3) leads to a low H<sub>2</sub>S content in the emission calculations. By using the highest concentration value of H<sub>2</sub>S given in Sonibare and Akeredolu (2004) (see Tab. 3, H<sub>2</sub>S concentration 0.03% instead of 0.005%), we try to estimate the upper limit of SO<sub>2</sub> emission, assuming that all flares are provided with this more sulfur containing gas. With this approach the

- 649 spatiotemporal<u>temporal</u> averaged <u>sum of</u>  $SO_2$  emissions <u>over SWA</u> increase from <u>0.636</u> to <u>4.9320</u> kg 650 h<sup>-1</sup>. The maximum values in the flare boxes increase from <u>4.7</u> to <u>41.8</u> kg h<sup>-1</sup>. These are rather low 651 values.
- This comparison reveals that among the flaring conditions also the natural gas composition plays an important role in estimating the flaring emissions reasonably. To rely on a single measurement dataset for a large flaring domain and without taking into account spatial variability is therefore problematic but has to be accepted owing to insufficient data shortage.
- This section has estimated the uncertainties in gas flaring due to cloud cover, parameters of IU14, the fraction of radiated heat, the temporal variability and the H<sub>2</sub>S concentration in the natural gas. The uncertainty regarding the spatial variability of the total hydrocarbon fraction of the natural gas, which is estimated by the variations in the ten flow station measurements of Sonibare and Akeredolu (2004), is below 1%.
- 661 However, there are further assumptions or sources of uncertainty which cannot be quantified within 662 this study: We assume that the natural gas composition, which is measured in one region, is valid for 663 SWA entirely. The gas flares are taken as constant emission sources because  $\frac{VNPVNF_{flare}}{VNPVNF_{flare}}$  only 664 provides one observation (overflight) per day. We cannot take into account the spatial variability of 665 the flares concerning the IU14 parameters and the stack heights. And finally IU14 delivers no VOCs 666 and black carbon.
- 667

# 668 **4. Comparison with existing emission inventories**

670 The following section places the estimated flaring emissions of this study in the context of existing 671 emission inventories, by taking the focus on CO<sub>2</sub>. A direct comparison with existing emission 672 inventories is problematic due to different reference time periods, spatial domains, definitions of 673 emission sectors and the limitation of chemical compounds. Tab. 5 summarizes the CO<sub>2</sub> emissions for different inventories regarding Nigeria or the Niger Delta area as denoted in Fig. 5a, as the flaring 674 675 hotspot of the research domain. The results of this study shows no flaring in the northern part of 676 Nigeria and therefore flaring within the Niger Delta area can be seen as the total flaring area of the 677 country. To derive annual emission values for the results of this study, it is assumed that the flaring 678 emission conditions of TP14 and TP15 are representative for the whole year 2014 and 2015 679 respectively. Therefore the hourly emissions are integrated over 365 days. In addition to the three 680 inventories E<sub>obs</sub>, E<sub>com</sub> and E<sub>clim</sub>, whose emissions are derived from temporal averages of the source 681 temperature and radiant heat, also an emission estimation using instantaneous source temperature 682 and radiant heat (calculating emissions for every single observation and subsequent temporal 683 averaging of the emissions) for both time periods is presented in Tab.5 (E<sub>clim</sub>, instant. input).

684

685**Tab.5.** Comparison between existing emission inventories for  $CO_2$  (with a focus on gas flaring if available) and the results of686this study for Nigeria or the Niger Delta area in teragram (Tg) per year. For TP14 and TP15 it is assumed that the two month687observations represent the flaring conditions of the whole year 2014 and 2015 respectively. Therefore the emissions were688integrated to yearly values. The values in brackets represent domain of uncertainty arising from the upperUP14 parameters689and lower limit owing to the uncertainties estimated in Section 3-spatial variability in total hydrocarbon is given in brackets.690For the fraction of radiated heat  $f_2$  the mean value 0.27 and the lower (upper) boundary of 0.16 (0.38) are used,691representing a further source of uncertainty. The products given in bold are directly related to flaring emissions.

Emission inventory	Time period	C	O <sub>2</sub> emissions (T	'g y⁻¹)
		f = 0.16	f = 0.27	f = 0.38
This study (E <sub>obs</sub> , averaged)	<u>2014 (from TP14)</u>	$\frac{1.7}{(2.2)}$	<u>1.0 (</u> <sup>1.3</sup> / <sub>0.2</sub> )	$0.7(^{1.0}/_{0.1})$
<u>This study (E<sub>com</sub>, averaged)</u>	<u>2014 (from TP14)</u>	$4.5(^{6.1}/_{0.9})$	$\frac{2.7}{(3.6)}$	$1.9(^{2.6}/_{0.3})$

This study (E <sub>clim</sub> )	2014 (from TP14)	<del>13 (<sup>19</sup>/3)</del> 4.9	<del>8 (<del>12</del>/<u>-</u>)2.9</del>	<del>6 (<sup>8</sup>/1)</del> 2.1
This study (E <sub>obs</sub> , averaged)	<u>2015 (from TP15)</u>	$\frac{(^{6.5}/_{1.0})}{1.0(^{1.4}/_{0.2})}$	( <sup>3.9</sup> / <sub>0.6</sub> ) <u>0.6 (<sup>0.8</sup>/<sub>0.1</sub>)</u>	$\frac{(^{2.8}/_{0.4})}{0.4(^{0.6}/_{0.0})}$
This study (E <sub>com</sub> , averaged)	2015 (from TP15)	$\frac{3.4}{(4.5)}$	$\frac{2.0}{(2.7)}$	$\frac{1.4}{(2.0)}$
This study (E <sub>clim</sub> )	2015 (from TP15)	<del>8 (<sup>12</sup>/<sub>2</sub>)3.7</del>	$\frac{5(7/1)2.2}{5}$	4 ( <del>5/)1.5</del>
		(4.9/0.7)	$(\frac{2.9}{0.4})$	$(^{2.1}/_{0.3})$
This study ( <del>E<sub>75</sub>E<sub>clim</sub>, instant.</del>	2014 (from TP14)	<del>7 (<sup>11</sup>/<sub>2</sub>)9.9</del>	4 <u>5.9</u>	<u>34.2</u>
<u>input</u> )		$(\frac{13.2}{2.0})$	( <del>7/1</del> 7.9/1.2)	( <del>5/</del> 5.6/8)
This study ( <del>E<sub>75</sub>E<sub>clim</sub>, instant.</del>	2015 (from TP15)	<del>5 (<sup>7</sup>/<sub>1</sub>)8.8</del>	<del>З (<sup>4</sup>/<sub>П</sub>)5.2</del>	<del>2 (<sup>3</sup>/<sub>Ω</sub>)3.7</del>
<u>input</u> )		$(\frac{11.8}{1.8})$	(7.0/1.0)	(4.9/0.7)
<b>CDIAC (2015b)</b> <sup>1</sup>	2011	1.0	27.47	
EIA (2015) <sup>2</sup>	2010; 2011; 2013		38.81; 41.39; 52	2.83
Doumbia et al. $(2014)^1$	2010		45	
<b>EDGAR 4.2</b> <sup>3</sup> (ECCAD, 2015)	2008		8.75	
<b>EDGAR 4.2</b> <sup>4</sup> (ECCAD, 2015)	2008		3.50	
EDGAR 4.3.2 <sup>5</sup> (EDGAR, 2016)	2010; 2011; 2012		29.4, 28.8, 28	.9
EDGARv43FT2012 <sup>6</sup> (EDGAR,	2014		93.87	
2014)				

<sup>1</sup>from gas flaring, Nigeria

694 695 <sup>2</sup>from consumption and flaring of natural gas

696 697 <sup>3</sup>from refineries and transformation, Nigeria

<sup>4</sup>from refineries and transformation, Niger Delta area according to Fig. 5a

698 699 <sup>5</sup>from venting and flaring of oil and gas production, Nigeria

<sup>6</sup>emission totals of fossil fuel use and industrial processes (cement production, carbonate use of limestone and dolomite, non-energy use of 700 fuels and other combustion). Excluded are: short-cycle biomass burning (such as agricultural waste burning) and large-scale biomass burning (such as forest fires), Nigeria

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The CO<sub>2</sub> emission estimations of this study are given in Tab.\_5 together with an overall uncertainty range (+38/\_72%)of (+33/\_79%) in brackets, including the uncertainty from the IU14 parameters  $\eta$  and  $\delta \left(\frac{+12}{-52} + \frac{12}{-53}\right)$  and the gauge pressure  $\left(\frac{+26}{-21} + \frac{20}{-25}\right)$  and from spatial variability 705 706 of total hydrocarbon. The latter uncertainty is small (below 1%) owing to the low variation in THC 707 concentration in the measurements of Sonibare and Akeredolu (2004). The uncertainty owing to the 708 fraction of radiated heat f is represented by using the average value of 0.27 and the upper and lower 709 estimate of 0.16 and 0.38 respectively. The uncertainty due to cloud cover is represented by E35-Regarding the relatively large uncertainty there is no preference the difference in one of the emission 710 711 estimates E<sub>clim</sub> and E<sub>obs</sub> and E<sub>com</sub>.

By assuming the uncertainty range of the fraction of radiated heat f between 0.16 and 0.38, the 712 713 results of the study on hand show CO2 emissions in the same order of magnitude as By assuming that 714 <u> $E_{com}$  with f = 0.27 represents the best emission estimate for this study and by integrating the above</u> mentioned sources of uncertainty, a total Nigerian CO<sub>2</sub> flaring emission of 2.7 ( $^{3.6}/_{0.5}$ ) Tg y<sup>-1</sup> for 2014 715 and 2.0  $(^{2.7}/_{0.4})$  Tg y<sup>-1</sup> for 2015 was derived. Due to the high uncertainties, the two estimates are not 716 717 statistically different. These values are one order of magnitude smaller than the values from the 718 Carbon Dioxide Information Analysis Center (CDIAC, 2015b), the Energy Information Administration (EIA, 2015) and the EDGARv.4.3.2 (EDGAR, 2016) database, with best results for f = 0.16 but with an 719 720 overall tendency to underestimate the emissions. Eclim shows smaller deviations to the existing 721 inventories than the cloud correction approach of E<sub>75</sub>. A direct comparison is hindered by a time lag of 3-4 years and missing information about the uncertainties of CDIAC. The values of EIA are higher 722 723 than those of CDIAC because EIA includes the consumption of natural gas in addition to gas flaring. 724 Doumbia et al. (2014) combines Defense Meteorological Satellite Program (DMSP) observations of

- flaring with the emission factor method to derive flaring emissions. The results agree with EIA (2015)but are 64% higher than CDIAC (2015b).
- The emission inventory EDGAR v4.2 (ECCAD, 2015) delivers 8.75 (3.50) Tg  $CO_2 y^{-1}$  for Nigeria (Niger Delta area) for the emission sector *refineries and transformation*, which is in good agreement with the results for the study on hand.
- As a benchmark for the flaring  $CO_2$ , the total  $CO_2$  emissions for Nigeria are given by EDGAR (2014),
- 731 (fossil fuel use and industrial processes). Taking EDGAR (2014) as a reference for total CO<sub>2</sub> emissions
- 732 of Nigeria, flaring emissions contributes by 9 ( $\frac{13}{2}$ )% (2014; E<sub>clim</sub> f = 0.27), 14 ( $\frac{20}{3}$ )% (2014; E<sub>clim</sub> f = 0.27)
- 733 f = 0.16 contribute with 2 ( $^{3.9}/_{0.0}$ )% (this study for 2014; E<sub>com</sub>), 9% (2008; ECCAD, 2015), 28% (2011;
- 734 CDIAC, 2015b), 48% (2010; Doumbia et al., 2014) or 56% (2013; EIA, 2015). The large spread between
- the different inventories emphasizes the large uncertainty within the estimation of emissions fromgas flaring.
- 737 By using the climatological approach with instantaneous source temperature and radiant heat input
- 738data ( $E_{clim}$ , instant. input) instead of temporal averages ( $E_{clim}$ ), the emissions are increased by approx.739a factor of two ( $5.9 (^{7.9}/_{1.2})$  Tg y<sup>-1</sup> for 2014,  $5.2 (^{7.0}/_{1.0})$  Tg y<sup>-1</sup> for 2015). This underlines that also the740preprocessing of the remote sensing data for the calculation of the emissions is a considerable741source of uncertainty. However, due to the high uncertainties also the two emission estimates with
- 742 and without instantaneous data are not statistically different.
- A shortcoming of the PEGASOS\_PBL-v2 (not shown) and the EDGAR v4.2 emission inventory is the lack of offshore flaring emissions in the Gulf of Guinea south of Nigeria. For CDIAC and EIA this cannot be verified since the data is only available as a single value per country.
- 746 The differences between the results of this study and the existing emission inventories might be 747 caused by an underestimation insufficient information about the efficiency of the flow rate by VNP 748 and Eq. 1combustion processes of SWA flares or by an inconsistent definition of emission source 749 sectors for the existing inventories.  $E_{clm}$ ,  $E_{75}E_{com}$ , Doumbia et al. (2014) and CDIAC (2015b) focus on 750 gas flaring, whereas other products also include natural gas consumption and emissions from 751 refineries and transformation, which also can include non-flaring emissions within and outside the 752 areas indicated as flaring area by the satellite imagery. In addition, the existing inventories do not 753 provide current values (time lag of 2 to 6 years) and therefore not consider the emission reduction 754 indicated by Fig. 9.
- 755

# 756 **5. Discussion and conclusions**757

- The gas flaring emission estimating method of Ismail and Umukoro (2014) (IU14) has been combined with the remote sensing flare location determination of the VIIRS Nightfire Prerun V2.1 Flares only (<u>VNPVNF</u>) (VIIRS, <u>20152015a</u>) for a new flaring emission parameterization. The parameterization combines equations of incomplete combustion with the gas flow rate derived from remote sensing parameters instead of using emission factors and delivers emissions of the chemical compounds CO, CO<sub>2</sub>, <u>SO<sub>2</sub></u>, NO and NO<sub>2</sub>.
- Within this study the parameterization was applied to southern West Africa (SWA) including Nigeria as the second biggest flaring country. Two two-month flaring observation datasets for June-/July 2014 (TP14) and June/July 2015 (TP15) were used to create a flaring climatology for both time periods. In this climatology all detected flares emit with their mean activity- (climatological approach).
- The uncertainties owing to missed flare observations by cloud cover, parameterization parameters, interannual variability and the natural gas compositions were assessed. It can be shown that the

- 771highest uncertainties arise from the IU14 parameters ( $^{+33}/_{-79}$ %), followed by the definition of the772fraction of radiated heat f and the IU14 parameters. The uncertainty arising from flares masked by773clouds is estimated as 61% on average in TP15.
- 774By using remote sensing the cloud cover observations, a correction factor for the flaring775climatological detection of VNF and by assuming that all cloud-covered flares are active, an776additional emission dataset was derived which reducescombines the mean emissions about 50%.777from the currently observed flares and the climatological emissions from cloud-covered (not778detected) flares (combined approach). These emissions are on average 9% smaller than the779climatology but 61% larger than the net observations.
- However, owing to the large uncertainty ranges, no significant difference between the climatological
   inventory and the cloud corrected combined inventory can be stated. Comparing the emissions of
   2014 and 2015, a reduction in the flaring area, density of active flares and a significant reduction in
   SWANigerian flaring emissions about 3025% can be observed, which underlines the need for more
   recent emission inventories.
- 785 The uncertainty due to the natural gas composition is compound dependent. The spatial variation in 786 total hydrocarbon is negligible but the availability of hydrogen sulfide, which exclusively determines 787 the amount of emitted SO<sub>2</sub>, cause large uncertainty. By taking the combustion efficiency to derive 788 the fraction of unburned natural gas, the amount of emitted VOCs might be estimated in addition to 789 the species of the study on hand but would also be linked to high uncertainties concerning the VOC 790 speciation. The uncertainty in VOC emission is increased drastically by natural gas which is vented 791 directly into the atmosphere instead of being flared, since the venting cannot be detected by 792 VNPVNF.
- With a focus on Nigeria, the CO<sub>2</sub> emission estimates of this study were compared with existing 793 inventories. For the <u>climatologycombined approach</u>, CO<sub>2</sub> emissions of  $\frac{8}{(12/2)^2 \cdot 7} \frac{(3.6/0.5)}{(0.5)}$  Tg y<sup>-1</sup> for 794 2014 and  $\frac{5}{(7/12.0)} (2.7/04)$  Tg y<sup>-1</sup> for 2015 were derived. EDGAR v4.2 for the year 2008 shows the 795 796 same order of magnitude when limiting to emissions from refineries and transformation. The results 797 of this study are one order of magnitude smaller compared to CDIAC (Carbon Dioxide Information 798 Analysis Center) is in the same order of magnitude as the results of this study.), Doumbia et al. (2014) 799 and EIA (Energy Information Administration) show emissions which are 2.4 and 2.8 times higher than 800 the results). This emission underestimation is not caused by an underestimation of the flared gas 801 volume. VNF<sub>flare</sub> includes an estimation of the annual sum of flared gas by country. For Nigeria the estimated values are 8.56 (7.64) bcm flared gas in 2014 (2015). Within this study- higher values of 802 803 37.89 (20.68) bcm for 2014 (2015) are derived.
- The deviations might be caused by uncertaintiesthe uncertainty in the flow rate derived by VNP 804 805 radiant heat, which can be assessed only rudimentary via the parameterefficiency of the fraction of 806 radiated heatflares concerning the combustion process and their operation. A lack of information 807 regarding the combustion efficiency together with the high sensitivity of the parameters within the 808 combustion equations of IU14 can lead to high uncertainties. Additionally, the usage of emission 809 factors in the existing inventories which did not take into account the spatiotemporal variability of 810 flaring, inconsistent emission sector definitions or the time lag of the emission inventories of 2-5 years can lead tocause deviations. The positive trend in Nigerian gas flaring CO<sub>2</sub> emissions derived by 811 EIA from 38.81 to 52.83 Tg y<sup>-1</sup> between 2010 and 2013 contradicts the findings of Doumbia et al. 812 (2014) and this study, which generally show a decrease in emissions from 1994 to 2010 and from 813 814 2014 to 2015, respectively. Based on the sensitivity study, which reveals high uncertainties of the 815 flaring emission, we conclude that there is no preference in the choice of one of the emission

- 816 estimates climatological and or the combined approach presented in this study. Therefore for simplicity we recommend the use of the climatological approach when using the R package. 817
- 818 Despite the generally large uncertainties in the estimation of emissions from gas flaring, this method
- 819 allows a flexible creation of flaring emission datasets for various applications (e.g. as emission 820 inventory for atmospheric models). It combines observations with physical based background
- 821
- concerning the combustion. The use of current data makes it possible to consider present trends in 822 gas flaring. Even the creation of near real-time datasets with a time lag of one day is possible. The
- 823 emissions are merged on grid predefined by the user and depending on the availability of VNPVNF
- 824 data, the temporal resolution can be selected from single days to years.
- 825 An improvement of this parameterization can be achieved by an extension of the IU14 method to 826 black carbon and VOCs and an inclusion of spatial resolved measurements of the natural gas 827 composition in combination with information of the gas flaring processes from the oil producing 828 industry.
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#### Code and/or data availability 841 842

- 843 This publication includes a package of well documented R scripts which is free available for research 844 purposes and enables the reader to create their own gas flaring emission datasets. It includes 845 exemplarily the preprocessing for June-/July 2015 with a focus on southern West Africa. You get 846 access to the code via zenodo.org (DOI: 10.5281/zenodo.50938).61151), entitled "Gas flaring 847 emission estimation parameterization v2".
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