Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions
Open Access

On computation of Hough functions

Houjun Wang[1,2], John P. Boyd[3], and Rashid A. Akmaev[2]
[1]CIRES, University of Colorado Boulder, Boulder, Colorado, USA.
[2]Space Weather Prediction Center, NOAA, Boulder, Colorado, USA.
[3]Department of Climate and Space Sciences and Engineering, University of Michigan, Ann Arbor, Michigan, USA.
Correspondence to: Houjun Wang (houjun.wang@noaa.gov)

Abstract.

Hough functions are the eigenfunctions of Laplace's tidal equation governing fluid motion on a rotating sphere with a resting basic state. Several numerical methods have been used in the past. In this paper, we compare two of those methods: *normalized* associated Legendre polynomial expansion and Chebyshev collocation. Both methods are not widely used, but both have some advantages over the commonly-used unnormalized associated Legendre polynomial expansion method. Comparable results are obtained using both methods. For the first method we note some details on numerical implementation. The Chebyshev collocation method was first used for Laplace tidal problem by Boyd (1976) and is relatively easy to use. A compact Matlab code is provided for this method. We also illustrate the importance and effect of including a *parity factor* in Chebyshev polynomial expansions for modes with *odd* zonal wavenumbers.

1 Introduction

Hough functions are the eigenfunctions of the eigenvalue problem of the following form:

$$\mathcal{F}(\Theta) + \gamma\Theta = 0, \tag{1}$$

where $\mathcal{F}$ is a linear differential operator, the *Laplace's tidal operator*, defined as:

$$\mathcal{F}(\Theta) \equiv \frac{d}{d\mu}\left(\frac{1-\mu^2}{\sigma^2-\mu^2}\frac{d\Theta}{d\mu}\right) - \frac{1}{\sigma^2-\mu^2}\left[\frac{s}{\sigma}\frac{\sigma^2+\mu^2}{\sigma^2-\mu^2} + \frac{s^2}{1-\mu^2}\right]\Theta, \tag{2}$$

with $\mu = \sin\phi \in [-1,1]$, ϕ the latitude, s the zonal wavenumber, and σ the dimensionless frequency normalized by 2Ω (Ω the earth's rotation rate), while

$$\gamma \equiv \frac{4a^2\Omega^2}{gh} \tag{3}$$

is the Lamb's parameter (Andrews et al., 1987, p. 154), with a the earth's radius, g the acceleration due to the earth's gravity, and h the so-called *equivalent depth*.

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions

Several numerical methods have been used to solve the eigenvalue problem for the Laplace tidal equation in the past. Hough (1898) pioneered the solutions of the Laplace tidal equations using spherical harmonic expansion, or equivalently *spherical harmonic Galerkin* method, so eigenfunctions of the eigenvalue problem Eq. (1) that describe the latitudinal dependence are often called *Hough functions* (Flattery, 1967; Longuet-Higgins, 1968; Lindzen and Chapman, 1969). The original method of computing Hough functions is based on expansion in terms of *unnormalized* associated Legendre polynomials (ALPs). Both Kato (1966) and Flattery (1967) used the *method of continued fractions* to solve for eigenvalues one by one with iterations. This is not the most convenient method to work with and some eigenvalues could be missed. Chen and Lu (2009) also discussed calculation of Hough functions following the same original formulation without showing any details on numerical procedures.

Computation of Hough functions based on expansion in terms of *normalized* ALPs was first used by Dikii (1965). It was later elaborated in a note by Groves (1981), along with a method of evaluating related wind functions. Jones (1970) used group-theoretical methods to obtain a matrix representation of Hough functions by expanding in normalized spherical harmonics.

Although it is closely related to the original method of expansion in terms of *unnormalized* ALPs, expansion in terms of the *normalized* ALPs leads to two symmetric matrices for symmetric and anti-symmetric modes. This has both *computational and conceptual* advantages over the original expansion in unnormalized ALPs: 1) the eigenvalue problem of symmetric matrix can be solved very accurately by Jacobi method (e.g., Demmel and Veselić, 1992), and 2) symmetry guarantees that all of the "eigenvalues are real and that there is an orthonormal basis of eigenvectors" (Golub and Van Loan, 1996, p. 393).

There is also another way of computing Hough functions or *global normal modes*, such as Longuet-Higgins (1968); Kasahara (1976); Žagar et al. (2015), also using spherical harmonic expansion, in which the equivalent depth is assigned (for each zonal wavenumber) and the frequency of the normal modes are obtained as the eigenvalues. This is different from eigenvalue problem for tidal waves in which the wave frequencies and zonal wavenumber are specified and eigenvalues are obtained and used to compute equivalent depths, just as stated in the original eigenvalue problem Eq. (1).

The Chebyshev collocation method was first used by Boyd (1976) to solve the eigenvalue problem for the Laplace tidal equation. It uses Chebyshev polynomials in the coordinate $\mu = \sin\phi$, which is equivalent to using an ordinary Fourier cosine or sine series in latitude. The Chebyshev collocation method is a general-purpose numerical method. Boyd (1976) listed several advantages of Chebyshev polynomial expansion over spherical harmonic expansion (basis function set becomes simpler and not restricted to spherical domain) as well as collocation method over Galerkin method (numerical quadrature is used to approximate the integrals). These advantages make it relative easy to work with Chebyshev collocation method than with spherical harmonic Galerkin method: derivation is no cumbersome and numerical implementation is straightforward. See also (Hesthaven et al., 2007, Chapter 3) for a discussion of advantages of Fourier-collocation methods over the Fourier-Galerkin methods.

A few remarks on unnormalized versus normalized ALP expansion are also in order here. The unnormalized polynomials (not just ALPs, but Legendre and Chebyshev and Hermite polynomials too) have survived because the canonical unnormalized forms have polynomial coefficients that are integers or rational numbers. This is convenient for many applications, such as when using exact arithmetic in computer algebra. Note that this property carries over to the Galerkin matrix elements for the Hough differential equation, which are rational functions of r and s in Eq. (6). Also, for some purposes it is very convenient

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions
Open Access
EGU

to use polynomials which are all 1 at x=1, as true for unnormalized Chebyshev and Legendre polynomials. The bad news is that unnormalized polynomials generate bigger roundoff errors in all calculations, not just computing matrix eigenvalues. The Galerkin matrix element formulas are more complicated for normalized polynomials. As we noted above, a particular advantage of working with normalized ALPs is that the discretization matrix becomes a symmetric matrix. Spectral discretizations often generate a few inaccurate eigenvalues with nonzero imaginary parts, but the eigenvalues of a symmetric tridiagonal matrix are always real.

In this paper we compare the solution of the eigenvalue problem for the Laplace tidal operator using two numerical methods, the normalized ALP expansion method and the Chebyshev collocation method. Both methods are not widely used, but both have some advantages over the commonly-used unnormalized ALP expansion. For the first method we note some details of numerical implementation as the denominators in some terms of matrix entries can become zero. For the second method a compact Matlab code is provided to facilitate its use. We also discuss other related issues and show that there is no accuracy penalty in using the Chebyshev collocation method.

2 Computation of Hough functions using normalized associated Legendre polynomial expansion

The first method uses the expansion in terms of *normalized* associated Legendre polynomials (ALPs) (e.g., Groves, 1981). To solve the Laplace's tidal equation, first expand Θ in terms of the *unnormalized* associated Legendre polynomials P_r^s

$$\Theta = \sum_{r=s}^{\infty} c_r P_r^s(\mu). \tag{4}$$

Substituting into the Laplace tidal equation Eq. (1), one obtains

$$Q_{r-2}c_{r-2} + (M_r - \lambda)c_r + S_{r+2}c_{r+2} = 0, \qquad (r \geqslant s), \tag{5}$$

where

$$Q_{r-2} = \frac{(r-s)(r-s-1)}{(2r-1)(2r-3)[s/\sigma - r(r-1)]}, \tag{6a}$$

$$\begin{aligned} M_r = {} & \frac{\sigma^2[r(r+1) - s/\sigma]}{r^2(r+1)^2} \\ & + \frac{(r+2)^2(r+s+1)(r-s+1)}{(r+1)^2(2r+3)(2r+1)[s/\sigma - (r+1)(r+2)]} \\ & + \frac{(r-1)^2(r^2-s^2)}{r^2(4r^2-1)[s/\sigma - r(r-1)]}, \end{aligned} \tag{6b}$$

$$S_{r+2} = \frac{(r+s+2)(r+s+1)}{(2r+3)(2r+5)[s/\sigma - (r+1)(r+2)]}, \tag{6c}$$

and

$$\lambda = \frac{gh}{4a^2\Omega^2} = \frac{1}{\gamma}. \tag{7}$$

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions Open Access EGU

These equations were first given by Hough (1898); see also Lindzen and Chapman (1969).

The *normalized* associated Legendre polynomials $P_{r,s}$ are defined in terms of the *unnormalized* associated Legendre polynomials P_r^s by

$$P_{r,s} = \left[\frac{2(r+s)!}{(2r+1)(r-s)!}\right]^{-\frac{1}{2}} P_r^s. \tag{8}$$

Expanding Θ in terms of the *normalized* associated Legendre polynomials $P_{r,s}$

$$\Theta = \sum_{r=s}^{\infty} a_r P_{r,s}(\mu), \tag{9}$$

we have (Dikii, 1965; Groves, 1981)

$$L_{r-2}a_{r-2} + (M_r - \lambda)a_r + L_r a_{r+2} = 0 \qquad (r \geqslant s), \tag{10}$$

where

$$L_r = \frac{[(r+s+1)(r+s+2)(r-s+1)(r-s+2)]^{\frac{1}{2}}}{(2r+3)[(2r+2)(2r+5)]^{\frac{1}{2}}[s/\sigma-(r+1)(r+2)]}, \tag{11a}$$

$$M_r = -\frac{\sigma^2-1}{(s/\sigma+r)(s/\sigma-r-1)} + \frac{(r-s)(r+s)(s/\sigma-r+1)}{(2r-1)(2r+1)(s/\sigma+r)[s/\sigma-r(r-1)]} + \frac{(r-s+1)(r+s+1)(s/\sigma+r+2)}{(2r+1)(2r+3)(s/\sigma-r-1)[s/\sigma-(r+1)(r+2)]}. \tag{11b}$$

Equation (10) can be written in a matrix form for the coefficients vector $x = [a_s, a_{s+1}, a_{s+2}, a_{s+3}, \ldots]^T$ as the matrix eigenvalue problem $F_0 x = \lambda x$, with matrix F_0 defined as

$$F_0 = \begin{bmatrix} M_s & 0 & L_s & 0 & 0 & \ldots \\ 0 & M_{s+1} & 0 & L_{s+1} & 0 & \ldots \\ L_s & 0 & M_{s+2} & 0 & L_{s+2} & \ldots \\ 0 & L_{s+1} & 0 & M_{s+3} & 0 & \ldots \\ 0 & 0 & L_{s+2} & 0 & M_{s+4} & \ldots \\ \vdots & \vdots & \vdots & \vdots & \vdots & \ddots \end{bmatrix}. \tag{12}$$

Or it may be written as, respectively, $F_1 x_1 = \lambda_1 x_1$, $x_1 = [a_s, a_{s+2}, \ldots]^T$ for symmetric modes, with matrix F_1 defined as

$$F_1 = \begin{bmatrix} M_s & L_s & 0 & 0 & \ldots \\ L_s & M_{s+2} & L_{s+2} & 0 & \ldots \\ 0 & L_{s+2} & M_{s+4} & L_{s+4} & \ldots \\ \vdots & \vdots & \vdots & \vdots & \ddots \end{bmatrix}, \tag{13}$$

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions Open Access EGU

and $F_2 x_2 = \lambda_2 x_2$, $x_2 = [a_{s+1}, a_{s+3}, \ldots]^T$ for antisymmetric modes, with matrix F_2 defined as

$$F_2 = \begin{bmatrix} M_{s+1} & L_{s+1} & 0 & 0 & \ldots \\ L_{s+1} & M_{s+3} & L_{s+3} & 0 & \ldots \\ 0 & L_{s+3} & M_{s+5} & L_{s+5} & \ldots \\ \vdots & \vdots & \vdots & \vdots & \ddots \end{bmatrix}. \tag{14}$$

These are real symmetric matrices and the eigenvalue problem can be solved accurately using the Jacobi methods (e.g., Golub and Van Loan, 1996, Chapter 8). The computed eigenvectors are the expansion coefficients.

A note on numerical implementation is relevant here, since denominators of terms in M_r can become zero. We found that form (6b), instead of the form of (11b), of M_r can be used to advantage, even though the two forms are *equivalent*. In addition, we should set that last term of (6b) of M_r to zero when it becomes a form of $0/0$. Thus, to compute the $(s = 2, \sigma = 1)$ modes or SW2 (*semidiurnal, westward propagating, zonal wave number 2*) modes, we should set the last term of (6b) to zero when $r = s = 2$.

The Fortran 90 source code of the Jacobi eigenvalue algorithm implemented by Burkardt (2013) can be used to solve the two symmetric matrix eigenvalue problems. It can actually, for the $(s = 1, \sigma = 0.5)$ modes or DW1 (*diurnal, westward propagating, zonal wave number 1*) tide, compute the one *infinity* eigenvalue with $P_{2,1}$ as the eigemode, "the most important odd mode" (Lindzen and Chapman, 1969, p. 151) since $P_{2,1} \propto \sin\phi\cos\phi$. So in this way we will not miss any important eigenvalue or eigenfunction; see Section 4 for a discussion on the "missing" modes for the solar diurnal modes and the completeness of Hough functions. When using Matlab, we can set any *inf* matrix entry to *realmax* and then use the Matlab function *eig* to solve the matrix eigenvalue problem. It is also preferable to compute eigenvalues for symmetric and anti-symmetric modes separately, especially when there are interior singularities, e.g., for the DW1 tide.

Using the method of expansions in the normalized associated Legendre polynomials, truncated at $r_{max} = 60$ on 94 Gaussian quadrature points, we compute eigenvalues and eigenfunctions for several important solar tides. We use *solar day* instead of *sidereal day* in our computations. The first several equatorial symmetric and anti-symmetric modes for DW1 are shown in Fig. 1. The first several equatorial symmetric and anti-symmetric modes for SW2 of scalar fields are shown in Fig. 2(a)-(b). The first several equatorial symmetric and anti-symmetric modes for $(s = 3, \sigma = 1.5)$ modes or TW3 (*terdiurnal, westward propagating, zonal wave number 3*) for temperature field are shown in Fig. 3. For completeness, a method of computing Hough functions for the horizontal wind components by Groves (1981) is presented in Appendix A.

3 Computation of Hough functions using Chebyshev collocation method

The Chebyshev collocation method was first used Boyd (1976) to solve Laplace tidal problem. Expand Θ in terms of the Chebyshev polynomials $T_n(\mu)$:

$$\Theta(\mu) = \sin^m \varphi \sum_{n=0}^{N} b_n T_n(\mu), \quad \text{with } m = mod(s,2), \tag{15}$$

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions

which includes a *parity factor* $\sin\varphi$ for the *odd* zonal wavenumber s (Orszag, 1974; Boyd, 1978), where φ is colatitude, $\varphi = \pi/2 - \phi$. Note that the Chebyshev collocation method uses Chebyshev polynomials in the coordinate of $\mu = \sin\phi$, which is equivalent to using an ordinary Fourier cosine or sine series in latitude, albeit on nonuniform distributed Chebyshev grids clustered near the two boundary points. Discussion on property of Chebyshev polynomials and collocation method can be found in Boyd (2001) and Trefethen (2000). A Matlab implementation is shown in Appendix B.

Parity requirement is discussed in Orszag (1974). To quote from Orszag (1974) "If parity requirements are violated, then differentiability is lost (at the boundaries, i.e., at the poles), possibly resulting in slow convergence of series expansions and associated Gibbs' phenomena. It is important that assumed spectral representations not impose an incorrect symmetry on a solution if infinite-order accurate results are desired" (see also Boyd (1978)).

To show how accuracy is affected by parity factor, we compare the eigenfunction expansion coefficients b_n computed with or without parity factor in Fig. 4. For both terdiurnal and pentadiurnal tides, when the parity factor is removed, only limited lower-order algebraic convergence rates are achieved: 4^{th}-order for terdiurnal and 7^{th}-order for pentadiurnal. When the parity factor is included, spectral or exponential convergence is restored. Thus including the parity factor improves the accuracy dramatically, so solutions are less affected by singularities when they exist. It is important to include the parity factor when computing eigenvalues and eigenfunctions for DW1 ($s = 1, \sigma = 0.5$) modes (see discussion below).

The Matlab code listed in Appendix B includes a *parity factor* for the odd zonal wavenumber. It also computes Hough modes for horizontal wind components. The computed eigenvalue in this case is just (negative) γ and from Eq (3) we can compute the corresponding equivalent depths h. Hough functions are simply the computed eigenvectors, with different normalization factors that are irrelevant, when Chebyshev differential matrices are used. So the eigenvalue and eigenvector problem we solve can be viewed as a direct discretization of the original operator eigenvalue problem (1).

Table 1 compares the number of good eigenvalues that can be obtained using the two methods. The "good" eigenvalue is defined as one whose *relative error*

$$E_{\text{rel}}(\hat{\lambda}) = \frac{|\lambda - \hat{\lambda}|}{|\lambda|}$$

is less than 10^{-6}, where λ is the eigenvalue computed at high truncation $N = 160$, considered to be accurate for purpose of comparison. It shows that for DW1 about 60% of the computed eigenvalues are good using the normalized ALP expansion method and about 50% of the computed eigenvalues are good using the Chebyshev collocation method; for SW2 a little over 50% of the computed eigenvalues are good using both methods; and for TW3 the number of good eigenvalues is about 75% for both methods. We note that for DW1 only about 15% of the computed eigenvalues are good *without parity factor*, contrasted to 50% *with parity factor*. This again illustrates the importance of preserving correct parity.

Considering the "unusual difficulties" in solving the eigenvalue problem of Laplace tidal equation using general numerical methods, as remarked by Bailey et al. (1991), it is *remarkable* that Chebyshev collocation method with a parity factor for odd zonal wavenumber can be used so successfully in solving the eigenvalue problem of the Laplace tidal equation.

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

4 A remark on the completeness of Hough functions

Although the completeness of Hough functions for zonal wavenumber s and period $T = (s+1)/2$ days was questioned earlier by Lindzen (1965), it was later proved by Holl (1970), see also Homer (1992). Giwa (1974) proved by direct computation that, for zonal wavenumber s and period $T = (s+1)/2$ days, Hough functions for tidal oscillations are the same as the associated Legendre polynomials P^s_{s+1} and Hough functions form a *complete* set of orthogonal functions.

One advantage in using the *normalized* associated Legendre polynomials as basis functions, as shown in Section 2, is that the eigenvalue problem becomes an eigenvalue problem for two real symmetric matrices, one for symmetric modes and one for anti-symmetric modes. The spectral theory of (Hermitian) symmetric matrices tells us that these real symmetric matrices have "a complete set of orthogonal eigenvectors, and that the corresponding eigenvalues are real" (e.g., Lax, 2002, Chapter 28). Thus this approach in a heuristic way shows the completeness of Hough functions.

5 Summary and Conclusions

In this paper, we briefly survey the numerical methods for computing eigenvalues and eigenvectors for the Laplace tidal operator. In particular we compare two numerical methods: the *normalized* associated Legendre polynomial (ALP) expansion and Chebyshev collocation. The *normalized* ALP expansion method leads to two symmetric matrices which can be solved very accurately. It also has an advantage in providing another conceptual understanding for the completeness of eigenfunctions (Hough functions) of Laplace tidal operator. We also note some details on numerical implementation.

The Chebyshev collocation method was first used by Boyd (1976) for computing the eigenvalues for the Laplace tidal problem. Here we compare this method with the ALP expansion and found that both are producing comparable results. Chebyshev collocation is a general-purpose numerical method and is relatively easy to work with. A compact Matlab code is provided to facilitate the use of Chebyshev collocation method for Laplace tidal problem.

The Chebyshev polynomial expansion method is merely a Fourier cosine expansion method in disguise (Boyd, 2001). In using the Chebyshev collocation method, it is important to include a *parity factor* in Chebyshev polynomial expansion for *odd* zonal wavenumber modes.

Appendix A: Hough functions for the horizontal wind components

Hough function for the horizontal wind components are (Groves, 1981; Lindzen and Chapman, 1969):

$$\Theta_u = \frac{(1-\mu^2)^{\frac{1}{2}}}{\sigma^2-\mu^2}\left[\frac{s}{1-\mu^2} - \frac{\mu}{\sigma}\frac{d}{d\mu}\right]\Theta, \tag{A1a}$$

$$\Theta_v = \frac{(1-\mu^2)^{\frac{1}{2}}}{\sigma^2-\mu^2}\left[\frac{(s/\sigma)\mu}{1-\mu^2} - \frac{d}{d\mu}\right]\Theta, \tag{A1b}$$

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions

for the eastward and northward components respectively. These can be evaluated numerically by discretizing the differential operators; or evaluated recursively as follows (Groves, 1981). Let

$$S_u = \cos\phi\,\Theta_u, \qquad S_v = \cos\phi\,\Theta_v, \tag{A2}$$

then from Eqs. (A1) we have

$$\sigma S_u - \mu S_v - (s/\sigma)\Theta = 0, \tag{A3a}$$

$$\mu S_u - \sigma S_v - (1/\sigma)\mathcal{D}\Theta = 0, \tag{A3b}$$

where $\mathcal{D} = (1-\mu^2)d/d\mu$. Note that there misses the factor of $1/\sigma$ before $\mathcal{D}\Theta$ in Eq. (40) of Groves (1981). For $s \geqslant 0$, we expand S_u and S_v in terms of the normalized associated Legendre polynomials:

$$S_u = \sum_{r=s}^{\infty} u_r P_{r,s}(\mu), \qquad S_v = \sum_{r=s}^{\infty} v_r P_{r,s}(\mu), \tag{A4}$$

and use Eq. (9) for expansions of Θ, as well as the recurrence relations for the normalized associated Legendre functions (which can be verified or derived from the recurrence relations for the unnormalized associated Legendre polynomials)

$$\mu P_{r,s} = b_r P_{r-1,s} + b_{r+1} P_{r+1,s}, \tag{A5a}$$

$$\mathcal{D} P_{r,s} = (r+1) b_r P_{r-1,s} - r b_r P_{r+1,s}, \tag{A5b}$$

where

$$b_r = [(r^2 - s^2)/(4r^2 - 1)]^{\frac{1}{2}}, \tag{A6}$$

then the coefficients of $P_{r-1,s}$ give

$$b_r u_r = \sigma v_{r-1} - b_{r-1} u_{r-2} - (1/\sigma)[(r-2)a_{r-2}b_{r-1} - (r+1)a_r b_r], \tag{A7a}$$

$$b_r v_r = \sigma u_{r-1} - b_{r-1} v_{r-2} - (s/\sigma) a_{r-1}. \tag{A7b}$$

The first several equatorial symmetric and anti-symmetric modes for SW2 ($s = 2, \sigma = 1$) for the zonal wind components computed using the above method are shown in Fig. 2(c)-(f). We also used the second-order central finite difference method to discretize the differential operators in Eqs. (A1a) and (A1b). Comparison of Hough mode computations for wind components using the method presented above and the finite difference method showing no visual differences, except at the two end points where the one-sided finite difference has to be used. The Matlab code for the Chebyshev collocation method also compute Hough functions for the horizontal wind components.

Appendix B: Listing of the Matlab codes for computing Hough functions

In this Appendix, we list the Matlab codes that can be used to compute eigenvalue and eigenvectors or Hough functions for the Laplace tidal equation. It includes a *parity factor* for modes with *odd* zonal wave number (s) (Orszag, 1974; Boyd, 1978).

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions
Open Access

```
% CHEB_HOUGH - Compute Hough functions
% using Chebyshev collocation methods
clear; format long e
a = 6.370d6; g = 9.81d0;
omega = 2.d0*pi/(24.d0*3600.d0);
%s = 1.d0; sigma = 0.4986348375d0; % DW1
 s = 1.d0; sigma = 0.5d0;     % DW1
%s = 2.d0; sigma = 1.0d0;     % SW2
%s = 3.d0; sigma = 1.5d0;     % TW3
parity_factor = mod(s,2);
N = 60; [D1,D2,x] = cheb_boyd(N,parity_factor);
a2 = (1-x.^2)./(sigma^2-x.^2);
a1 = 2.*x.*(1-sigma^2)./(sigma^2-x.^2).^2;
a0 = -1./(sigma^2-x.^2).*((s/sigma) ...
     .*(sigma^2+x.^2)./(sigma^2-x.^2) ...
     +s^2./(1-x.^2));
A = diag(a2)*D2 + diag(a1)*D1 + diag(a0);
[V,D] = eig(A); lamb = real(diag(D));
% sort eigenvalues and -vectors
[foo,ii] = sort(-lamb);
lamb = lamb(ii); hough = V(:,ii);
% equivalent depth (km)
h = -4.d0*a^2*omega^2/g./lamb/1000.d0;
% compute Hough functions for wind components
b1 = (sigma^2-x.^2).*sqrt(1.d0-x.^2);
b2 = sqrt(1.d0-x.^2)./(sigma^2-x.^2);
hough_u = diag(s./b1)*hough ...
          - diag(b2.*x./sigma)*D1*hough;
hough_v = diag((s/sigma).*x./b1)*hough ...
          - diag(b2)*D1*hough;
clf % plot Hough functions
for j = 1:60
u = hough(:,j); subplot(10,6,j)
plot(x,u,'.','markersize',8), grid on
xx = -1:.01:1; uu = polyval(polyfit(x,u,N),xx);
line(xx,uu)%, axis off
end
```

And here is the list of the Matlab codes for computing Chebyshev differential matrices *numerically* with an option for including the parity factor.

```
function [D1, D2, x] = cheb_boyd(N, pf)
% CHEB_BOYD - Compute differential matrix
```

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016


```
% for Chebyshev collocation method;
% It contains an optional parity factor (pf)
t = (pi/(2*N)*(1:2:(2*N-1)))';
x = cos(t); n = 0:(N-1);
ss = sin(t); cc = cos(t);
sx = repmat(ss,1,N); cx = repmat(cc,1,N);
nx = repmat(n,N,1);  tx = repmat(t,1,N);
tn = cos(nx.*tx);
if  pf==0
    phi2 = tn;
    PT = -nx.*sin(nx.*tx);
    phiD2 = -PT./sx;
    PTT = -nx.^2.*tn;
    phiDD2 = (sx.*PTT-cx.*PT)./sx.^3;
else
    phi2 = tn.*sx;
    PT = -nx.*sin(nx.*tx).*sx + tn.*cx;
    phiD2 = -PT./sx;
    PTT = -nx.^2.*tn.*sx ...
          - 2*nx.*sin(nx.*tx).*cx - tn.*sx;
    phiDD2 = (sx.*PTT-cx.*PT)./sx.^3;
end
D1 = phiD2 /phi2; % the first derivatives
D2 = phiDD2/phi2; % the second derivatives
```

References

Andrews, D., Holton, J., and Leovy, C.: Middle Atmosphere Dynamics, Academic Press, Inc., New York, 1987.

Bailey, P., Everitt, W., and Zettl, A.: Computing eigenvalues of singular Sturm-Liouville problems, Results Math., 20, 391–423, doi:10.1007/BF03323182, http://link.springer.com/10.1007/BF03323182, 1991.

Boyd, J. P.: Planetary waves and the semiannual wind oscillation in the tropical upper stratosphere, Ph.d. thesis, Harvard University, 1976.

Boyd, J. P.: The choice of spectral functions on a sphere for boundary and eigenvalue problems: A comparison of Chebyshev, Fourier and Associated Legendre expansions, Mon. Weather Rev., 106, 1184–1191, 1978.

Boyd, J. P.: Chebyshev and Fourier Spectral Methods, Dover Publications, Inc., 2 edn., 2001.

Burkardt, J.: JACOBI_EIGENVALUE: Eigenvalues and Eigenvectors of a Symmetric Matrix, http://people.sc.fsu.edu/~jburkardt/f_src/jacobi_eigenvalue/jacobi_eigenvalue.html, 2013.

Chen, Z.-Y. and Lu, D.-R.: On the Calculation of Hough Functions for Resolving Atmospheric Thermal Tidal Structure, Chinese J. Geophys., 52, 547–554, http://manu16.magtech.com.cn/geophy_en/EN/abstract/abstract1053.shtml, 2009.

Demmel, J. and Veselić, K.: Jacobi's method is nore accurate than QR, SIAM J. Matrix Anal. Appl., 13, 1204–1245, doi:10.1137/0613074, http://epubs.siam.org/doi/abs/10.1137/0613074, 1992.

Dikii, L. A.: The terrestrial atmosphere as an oscillating system, Izv. Atmos. Ocean. Phys. - English Ed., 1, 275–286, 1965.

Flattery, T. W.: Hough Functions, Tech. rep., University of Chicago, Chicago, 1967.

Giwa, F. B. A.: The 'missing' member of Hough's functions in diurnal tides, Pure Appl. Geophys., 112, 519–522, doi:10.1007/BF00877288, http://link.springer.com/10.1007/BF00877288, 1974.

Golub, G. H. and Van Loan, C. F.: Matrix Computations, The John Hopkins University Press, Baltimore and London, 3 edn., 1996.

Groves, G.: Notes on obtaining the eigenvalues of Laplace's tidal equation, Planet. Space Sci., 29, 1339–1344, doi:10.1016/0032-0633(81)90100-8, http://linkinghub.elsevier.com/retrieve/pii/0032063381901008, 1981.

Hesthaven, J. S., Gottlieb, S., and Gottlieb, D.: Spectral Methods for Time-Dependent Problems, Cambridge University Press, 2007.

Holl, P.: The completeness of the orthogonal system of the Hough functions, Nachrichten der Akad. der Wissenschaften Gottingen, 7, 159–168, 1970.

Homer, M. S.: Spectral Properties of the Laplace Tidal Wave Equation, J. London Math. Soc., s2-45, 79–96, doi:10.1112/jlms/s2-45.1.79, 1992.

Hough, S. S.: On the application of harmonic analysis to the dynamical theory of the tides. Part II: On the general integration of Laplace's dynamical equations, Philos. Trans. R. Soc. London, A191, 139–185, http://www.jstor.org/stable/90743, 1898.

Jones, M. N.: Atmospheric Oscillations-I, Planet. Space Sci., 18, 1393–1416, 1970.

Kasahara, A.: Normal modes of ultralong waves in the atmosphere, Mon. Weather Rev., 104, 669–690, http://journals.ametsoc.org/doi/abs/10.1175/1520-0493(1976)104%3C0669:NMOUWI%3E2.0.CO%3B2, 1976.

Kato, S.: Diurnal Atmospheric Oscillation 1. Eigenvalues and Hough Functions, J. Geophys. Res., 71, http://onlinelibrary.wiley.com/doi/10.1029/JZ071i013p03201/full, 1966.

Lax, P. D.: Functional Analysis, Wiley-Interscience, 2002.

Lindzen, R.: On the Asymmetric Diurnal Tide, Pure Appl. Geophys., pp. 142–147, http://link.springer.com/article/10.1007/BF00875296, 1965.

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Lindzen, R. S. and Chapman, S.: Atmospheric Tides, Space Sci. Rev., 10, 3–188, http://link.springer.com/content/pdf/10.1007/978-94-010-3399-2.pdf, 1969.

Longuet-Higgins, M.: The eigenfunctions of Laplace's tidal equations over a sphere, Philos. Trans. R. Soc. London, 262, 511–607, http://www.jstor.org/stable/73582, 1968.

Orszag, S. a.: Fourier Series on Spheres, Mon. Weather Rev., 102, 56–75, doi:10.1175/1520-0493(1974)102<0056:FSOS>2.0.CO;2, http://journals.ametsoc.org/doi/abs/10.1175/1520-0493(1974)102<0056:FSOS>2.0.CO;2, 1974.

Trefethen, L. N.: Spectral Methods in Matlab, SIAM: Society for Industrial and Applied Mathemetics, Philadelphia, PA, 2000.

Žagar, N., Kasahara, a., Terasaki, K., Tribbia, J., and Tanaka, H.: Normal-mode function representation of global 3-D data sets: open-access software for the atmospheric research community, Geosci. Model Dev., 8, 1169–1195, doi:10.5194/gmd-8-1169-2015, http://www.geosci-model-dev.net/8/1169/2015/, 2015.

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions
Open Access

Table 1. Number of good eigenvalues of three tidal waves DW1, SW2 and TW3 computed with different trunction N using two different methods: I - normalized ALP expansion, II - Chebyshev collocation.

N	DW1-I	DW1-II	SW2-I	SW2-II	TW3-I	TW3-II
8	2	0	2	0	3	1
16	6	1	6	5	10	6
24	10	3	10	9	16	13
32	16	9	14	13	22	19
40	22	14	20	18	28	25
48	28	15	24	22	36	32
56	32	24	29	27	42	39
64	38	29	34	32	48	45
72	43	29	38	37	56	52
80	49	39	44	42	62	59

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

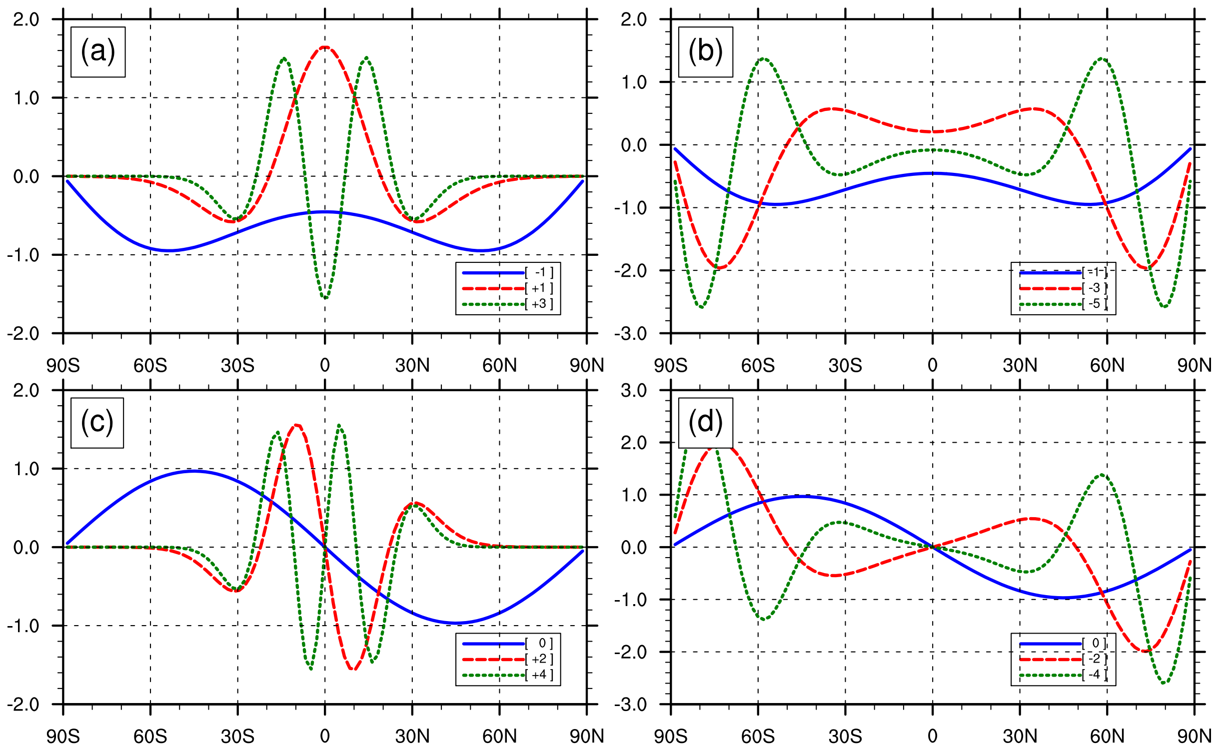

Figure 1. The first few symmetric and antisymmetric Hough modes for DW1 ($s = 1, \sigma = 0.5$) of scalar fields, computed using the normalized associated Legendre polynomial (ALP) expansions. Panels (a) and (b) are for symmetric modes, (c) and (d) are for anti-symmetric modes. The labels are: [-1] for the first *negative* mode with largest *negative* eigenvalue, [+1] for the first *positive* mode with largest *positive* eigenvalue, and [0] for the so-called missing mode with *zero* eigenvalue or *infinite* equivalent depth.

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development
Open Access
Discussions

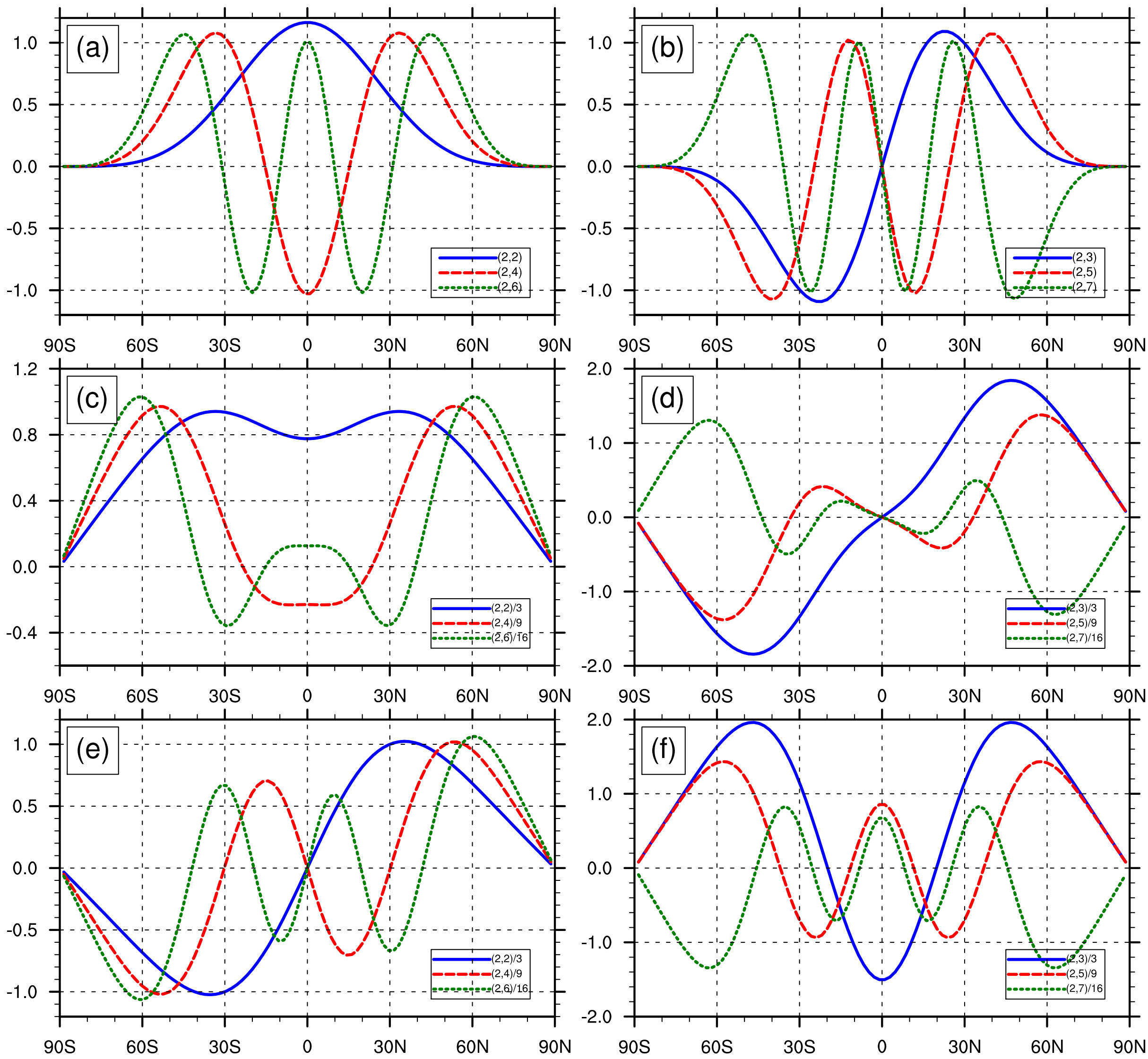

Figure 2. The first few symmetric and antisymmetric Hough modes for SW2 ($s = 2, \sigma = 1$), computed using the normalized associated Legendre polynomial (ALP) expansions. The left panels are symmetric modes and the right panels are anti-symmetric modes, except panels (e) and (f) which are reversed. Panels (a) and (b) are for the scalar fields, (c) and (d) for the zonal wind component, (e) and (f) for the meridional wind component. The labels are conventional.

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development Discussions
Open Access

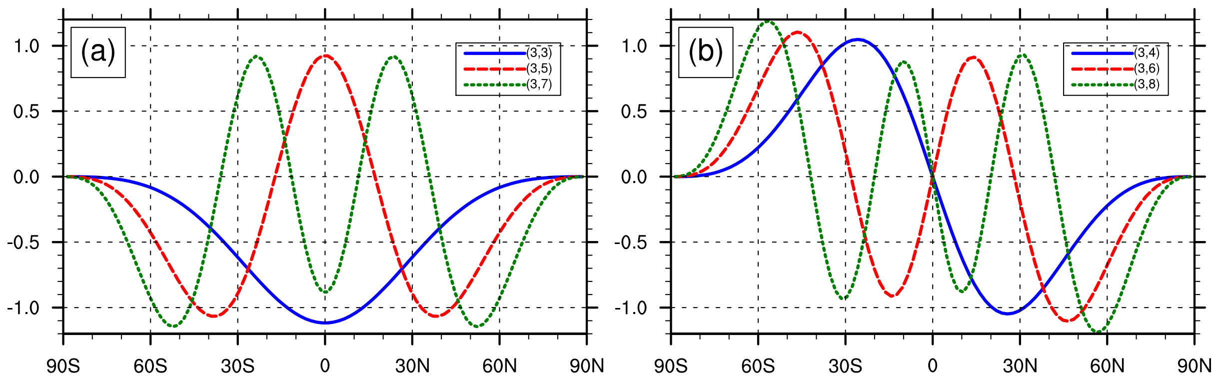

Figure 3. The first few symmetric and antisymmetric Hough modes for TW3 ($s = 3, \sigma = 1.5$) of scalar fields, computed using the normalized associated Legendre polynomial (ALP) expansions. The left panels are symmetric modes and the right panels are anti-symmetric modes.

Geosci. Model Dev. Discuss., doi:10.5194/gmd-2015-282, 2016
Manuscript under review for journal Geosci. Model Dev.
Published: 19 January 2016

Geoscientific Model Development
Open Access
Discussions

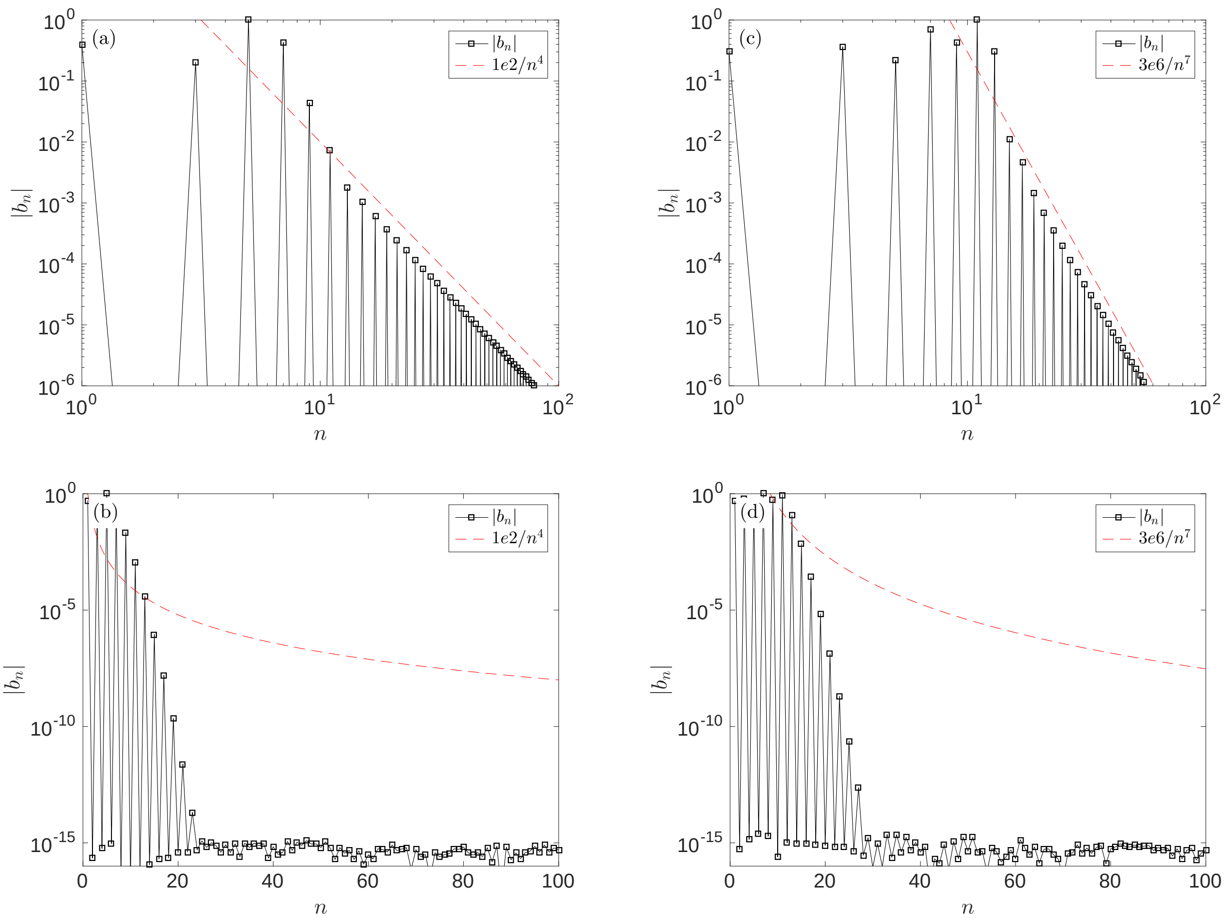

Figure 4. The absolute value of the expansion coefficients b_n in Eq. (15), truncated at $N = 150$. The left panels are for the terdiurnal tides, s=3, σ=1.5, for eigenfunction with eigenvalue γ=17.2098: (a) without parity factor, (b) with parity factor; The right panels are for penta-diurnal tides s=5, σ=2.5, for eigenfunction with eigenvalue γ=22.9721: (c) without parity factor, (d) with parity factor. An empirical fitting curve is also shown in red dash.