

We thank the two anonymous referees for their valuable comments and constructive suggestions on the manuscript. Below, we explain how the comments and suggestions are addressed and make note of the revision we made in the manuscript.

Anonymous Referee #1

General comments:

- *This study investigates the impact of different land surface parameterizations and vegetation distributions on emissions and mixing ratios of biogenic VOCs (and related oxidation products) simulated in California. Isoprene, MACR, MVK and monoterpenes are especially considered. Two different versions of the Model of Emissions of Gases and Aerosols from Nature (MEGAN v2.0 and MEGAN v2.1), together with two different land surface schemes (Noah and CLM4.0) and 5 different vegetation distributions (VEGM, USGS, VEG1, VEG2, VEG3) are alternatively used. Data collected during two field campaigns, CalNex and CARES, providing ground-based or flight observations, are also considered for model evaluation.*

This manuscript is well written and clearly presents an extensive work, which I really enjoyed reading, a work that provides clues to better understand the variability and uncertainty of biogenic VOC estimates between models. To some extent, the manuscript lacks of precise information, especially regarding the model framework. For example, the differences in emission calculation between the two versions of MEGAN used, or the connexions between the emission model and the land-surface scheme should be better described, in order to fully understand the possible source of variability in results provided. I therefore give a list of corrections and comments to improve the clarity of the manuscript, which I warmly recommend for publication in GMD.

We thank the reviewer for a detail review. Both text and figures are revised as the reviewer suggested.

Specific comments:

- *Section 2.2 and 2.3:*

These sections lack of clear information regarding the differences between the emission models or land-surface schemes, and connexions between them. First if CLM4 considers 16 PFTs, how many are taken into account in Noah?

Noah uses the 24 USGS land-use types. We have modified the text to include “Noah has four soil layers, with a total depth of two meters and a single slab snow layer that is lumped with the top-soil layer, which is set to a combined depth of 10 cm. It uses the 24 United States Geological Survey (USGS) land-use types, and does not treat sub-grid scale variability within a model grid cell.”

More clarification is added into Section 2.2 and 2.3 as indicated in the responses to the comments below.

- *From page 9, lines 199-203, it is not really clear to me which meteorology is considered when using MEGAN v2.0: is it eventually provided by WRF-CHEM or*

based on a monthly climatology?

MEGAN v2.0 in WRF-Chem needs instantaneous and past-days' mean meteorological variables that are from the WRF-Chem simulation and the monthly climatology dataset, respectively. We have modified the text to state “The biogenic emission calculation in MEGAN uses both instantaneous and the past-days' surface air temperature and solar radiation. MEGAN v2.0 obtains the instantaneous value from the land surface parameterization and the past-days' value from the climatological monthly mean dataset. In contrast, MEGAN v2.1 obtains both values directly from CLM.”

- *Differences in emission schemes between MEGAN 2.0 and MEGAN 2.1 should also be more precisely stated in the text regarding number of vegetation classes, emission factors (are they prescribed for each PFT for both MEGAN v2.0 and MEGAN v2.1 or is one using EF maps?).*

The text has been modified to include “In this study, both MEGAN v2.0 and v2.1 estimate biogenic species emissions based on the PFT distributions and the PFT specific emission factors. MEGAN v2.0 and v2.1 use 4 and 16 PFTs, respectively, as described below in Section 2.4.”

- *Connexions and variables coupling between emission model and land-surface scheme (any version) should be given in details: which of the variables calculated by the landsurface scheme are actually used in MEGAN v2.0 and v2.1 to calculated emissions? This is also especially important in section 4, when analyzing the impact of using different land-surface parameterizations.*

We have modified the text to say “In the released version, MEGAN v2.0 can be used with any land surface scheme available in WRF-Chem including Noah and CLM4.” and “The biogenic emission calculation in MEGAN uses both instantaneous and the past-days' surface air temperature and solar radiation. MEGAN v2.0 obtains the instantaneous value from the land surface parameterization and the past-days' value from the climatological monthly mean dataset. In contrast, MEGAN v2.1 obtains both values directly from CLM.” In section 4, we have revised the discussion in the following text: “Although the two land surface parameterizations produce slightly different values of surface temperature (Fig. 1), soil moisture (not shown), and net solar radiation near the surface (not shown), their impact on the biogenic emissions was small.” and “Although both experiments with Noah and CLM4 (red and orange lines, respectively) simulate similar isoprene emission fluxes with the maximum in the afternoon (Fig. 10), their respective isoprene+MVK+MACR mixing ratios are different at the four sites, particularly at site T0, where the Mv20CLM simulated isoprene+MVK+MACR mixing ratios during the daytime are about a factor of 2 larger than those from Mv20Noah. This inconsistency mainly results from the differences in the near surface meteorology, such as net surface radiation and temperature, between the two experiments (not shown) that

affects photochemistry, but this impact of surface meteorology occurs only at limited locations.”

- *Finally, nothing is said anywhere in the manuscript about the leaf area index, which is yet a crucial driving variable in emission estimate in MEGAN. How is it taken into account: is LAI prescribed or is it calculated by each land-surface scheme, and if so what are the LAI differences or similarities between them?*

The leaf area index is prescribed using the 4 PFTs in MEGAN2.0 and 16 PFTs in MEGAN2.1. Figure 6 has been added to show the difference in LAI among the experiments and the following description has been added to the text: “Figure 6 shows the spatial distributions of LAI used in MEGAN v2.0 and v2.1. The differences in the spatial distributions of LAI can significantly affect the biogenic emission calculation in MEGAN. It should be noted that in MEGAN v2.0 used in WRF-Chem, the LAI used for the calculation of the biogenic emissions is prescribed using the 4 PFTs, which is different than the land scheme that uses the LAI derived from the 24 USGS land categories.”

- *Page 9, lines 186-194: please also specify here in the text that MEGAN v.2.0 considers 4 PFTs only.*

Done, we have revised the text to say “MEGAN v2.0 and v2.1 use 4 and 16 PFTs, respectively, as described below in Section 2.4.”

- *Results from both MEGAN v2.0 and v2.1 are eventually compared with each other, and with observations. Is this comparison actually consistent since MEGAN v2.0 emission factors represent the net emission flux into the atmosphere, and MEGAN v2.1 ones the net primary emission that escape into the atmosphere? Are there significant differences between the two set of emission factors? MEGAN v2.0 emission factors should also be given, as is done for MEGAN v2.1 in figure 3. Ideally, maps of emission factors, projecting emission factor values over PFT distribution, would really help understanding the differences between both emission models.*

The difference in the definition (net flux versus primary emission) of emission factors affects the emission factors of compounds with bidirectional exchange but does not impact MEGAN isoprene and monoterpene emission factors because they have small deposition rates relative to emission rates. We have revised the text to state “The difference in the definition (net flux versus primary emission) of emission factors affects the emission factors of compounds with bidirectional exchange but does not impact MEGAN isoprene and monoterpene emission factors because they have small deposition rates relative to emission rates.”

Figure 4 has been revised to include the biogenic isoprene emission factor for the 4 PFTs used in MEGAN v2.0. Figure 5 has been added to show the differences in the spatial distributions of averaged biogenic isoprene emission factor in MEGAN v2.0 and v2.1 with different PFTs. The text is revised as “**Figure 4 shows the biogenic isoprene emission factor for each PFT prescribed in MEGAN v2.0 and MEGAN v2.1 in CLM4. In MEGAN v2.1, it shows that temperate broadleaf deciduous tree (PFT 7 listed in Table 1) has a large isoprene emission factor, while temperate needleleaf evergreen tree (PFT 1 listed in Table 1) has a small isoprene emission factor. A similar difference between broadleaf trees and needleleaf trees is indicated for MEGAN v2.0. Figure 5 shows the spatial distributions of averaged biogenic isoprene emission factor used in MEGAN v2.0 and v2.1 with different PFTs. It is evident that the difference in the distributions of PFTs results in a significant difference in spatial distributions of the isoprene emission factor.**”

Technical corrections:

- ***Page 3, line 69; page 4, line 74; page 5, line 105: change “BVOCs” to “BVOC”***
Corrected.
- ***Page 5, line 100: change “during the day but a factor of three” to “during the day but by a factor of three”***
Corrected.
- ***Page 7, line 145: please write what RRTMG stands for***
Corrected.
- ***Page 8, line 172: change “PFT’s to “PFTs”***
Corrected.
- ***Page 9, line 191: change “defined” to “defines”***
Corrected.
- ***Page 10, line 215: change “MEGAN to “MEGAN v2.1***
Corrected.
- ***Page 12, line 276: change “PFT’s” to “PFTs”***
Corrected.
- ***Page 14, line 312: change “BVOCs simulation” to “BVOC simulation”***
Corrected.
- ***Page 18, line 396: change “monterpene” to “monoterpene”***

Corrected.

- *Page 19, line 422: change “and monoterpene” to “and monoterpenes”*

Corrected.

- *Page 19, line 423: add “and Figure 13” (for monoterpenes) after “Figure 12” Page*

Corrected.

- *21, line 463-464: change “while both experiments are slightly smaller” to “while both experiment mixing ratios are slightly smaller”*

Corrected.

- *Page 28, line 634: change “and hence the atmospheric VOC mixing ratios” to “and hence of atmospheric VOC mixing ratios”*

Corrected.

- *Page 30, lines 688, 689 and 690: change “BVOCs emission” to “BVOC emission”*

Corrected.

- *Page 31, line 701: change “v20” and “v21” (twice) to “v2.0” and “v2.1” respectively*

Corrected.

- *Page 41, line 915: the font used for “Müller J.” seems different to me than the one used for the rest of the text*

Corrected.

- *Table 1 and Figure 2 captions: change “PFT’s” to “PFTs”*

Corrected.

- *Figure 12, bottom left plot: Is actually isoprene mixing ratio plotted or isoprene+MVK+MACR?*

It is isoprene. Now it is clarified as “At the Bakersfield site, only isoprene mixing ratios were reported so that the comparison is for isoprene only.”

Anonymous Referee #2

General comments:

- *This is an excellent paper and it should make a significant contribution to GMD. Because it will serve as a reference for users of the widely used community modeling system WRF-Chem, I agree with comments from reviewer #1 that it requires a bit more information and precision. In addition to the comments from the other reviewer, I would like to see more details and clarification on the following points.*

We thank the reviewer for a detailed review. Both text and figures are revised as the reviewer suggested.

Specific comments:

- *(1) The authors apply nudging. While is appropriate for their application in which they only look at the sensitivity of biogenic emissions to land surface parameterizations and vegetation distributions, the reader should have a little bit more info. Is the nudging also applied in the Boundary Layer (BL) and at the surface? Why did you choose not to nudge moisture? I would not expect the answers to this question to alter the quality of the results.*

The nudging is only applied in the free atmosphere above the BL. In general, we do not nudge moisture because we want the model to freely simulate clouds. As the reviewer also points out, the nudging method should not lead to changes of our results. It is now clarified in the text as “**The modeled u and v wind components and temperature in the free atmosphere above the planetary boundary layer are nudged towards the NARR reanalysis data with a time scale of 6 hours [Stauffer and Seaman, 1990].**”

- *(2) I would have been interested to get a bit more info on the difference in surface meteorology, assuming that nudging was not applied in the BL. What was the relative impact from meteorology compared to land-use and/or a different version of MEGAN? Of course, if nudging was applied in the BL this would be a moot point. If the authors can elaborate a little on this that could be useful.*

In this manuscript we have not focused on the meteorological impact. There are small differences in the surface meteorological fields among the experiments, for example, there are differences in latent heat and sensible heat fluxes. However, as we discussed in the text that the impact of surface meteorological difference on biogenic emissions is relatively small compared to the vegetation impact. For example in the text, “**Although the two land surface parameterizations produce slightly different values of surface temperature (Fig. 1), soil moisture (not shown), and net solar radiation near the surface (not shown), their impact on the biogenic emissions was small.**”

We also discussed about the potential impact of surface meteorology on surface mixing ratios. The text as been modified as follows, “**Although both experiments with Noah and**

CLM4 (red and orange lines, respectively) simulate similar isoprene emission fluxes with the maximum in the afternoon (Fig. 10), their respective isoprene+MVK+MACR mixing ratios are different at the four sites, particularly at site T0, where the Mv20CLM simulated isoprene+MVK+MACR mixing ratios during the daytime are about a factor of 2 larger than those from Mv20Noah. This inconsistency mainly results from the differences in the near surface meteorology, such as net surface radiation and temperature, between the two experiments (not shown) that affects photochemistry, but this impact of surface meteorology occurs only at limited locations.”

- *(3) I assume this was a dry period in the model simulations, so slight differences in cloud distributions could not have contributed much to the differences between model simulations in this case. However, could this have played a role in under/over forecasting for simulations of all runs in general?*

Yes, this study during a dry and warm period that favors biogenic emissions. For a more general case, the absolute impact may be smaller. The more quantitative conclusion should be drawn with multiple-season simulation in future studies. This is now acknowledged in the discussion section, where the text has been modified to read “It is also noteworthy that this study is in a relatively dry and warm season; therefore the impact of biogenic emission treatments may change for other seasons and during periods with higher cloudiness. A multiple-season investigation may be needed in future.”

1 **Sensitivity of biogenic volatile organic compounds (BVOCs) to land**
2 **surface parameterizations and vegetation distributions in California**

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19 Manuscript for submission to *WRF-Chem special issue in Geosci. Model Dev.*

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24 **Abstract.** Current climate models still have large uncertainties in estimating biogenic
25 trace gases, which can significantly affect atmospheric chemistry and secondary aerosol
26 formation that ultimately influences air quality and aerosol radiative forcing. These
27 uncertainties result from many factors, including uncertainties in land-surface processes
28 and specification of vegetation types, both of which can affect the simulated near-surface
29 fluxes of biogenic volatile organic compounds (BVOCs). In this study, the latest version
30 of Model of Emissions of Gases and Aerosols from Nature (MEGAN v2.1) is coupled
31 | within the land surface [scheme](#) CLM4 in the Weather Research and Forecasting model
32 | with chemistry (WRF-Chem). In this implementation, MEGAN v2.1 shares a consistent
33 | vegetation map with CLM4 for estimating BVOC emissions. This is unlike MEGAN v2.0
34 | in the public version of WRF-Chem that uses a standalone vegetation map that differs
35 | from what is used by land surface [schemes](#). This improved modeling framework is used
36 | to investigate the impact of two land surface [schemes](#), CLM4 and Noah, on BVOCs and
37 | examine the sensitivity of BVOCs to vegetation distributions in California. The
38 | measurements collected during the Carbonaceous Aerosol and Radiative Effects Study
39 | (CARES) and the California Nexus of Air Quality and Climate Experiment (CalNex)
40 | conducted in June of 2010 provide an opportunity to evaluate the simulated BVOCs.
41 | Sensitivity experiments show that land surface [schemes](#) do influence the simulated
42 | BVOCs, but the impact is much smaller than that of vegetation distributions. This study
43 | indicates that more effort is needed to obtain the most appropriate and accurate land
44 | cover datasets for climate and air quality models in terms of simulating BVOCs, oxidant
45 | chemistry, and consequently secondary organic aerosol formation.

46 1. Introduction

47 Volatile organic compounds (VOCs) in the atmosphere play an important role in
48 atmospheric chemistry, and therefore can significantly affect ozone and secondary
49 organic aerosol (SOA) formation and ultimately air quality and climate [e.g., Chameides
50 et al., 1992; Fehsenfeld et al., 1992; Andreae and Crutzen, 1997; Pierce et al., 1998;
51 Poisson et al., 2000; Sanderson et al., 2003; Claeys et al., 2004; Arneth et al., 2010].
52 Significant effort has been made on obtaining accurate predictions of atmospheric VOC
53 concentrations; however, there remain large differences between observed and simulated
54 values. These uncertainties result from many factors, including biogenic emission rates
55 that are influenced by near-surface meteorological processes, sub-surface processes,
56 representation of vegetation distributions, and plant biology [Guenther et al., 2013].

57 Biogenic emissions are a major source of VOCs [e.g., Zimmerman et al., 1978;
58 Mueller, 1992] in the atmosphere. In particular, isoprenoids (consisting mainly of
59 isoprene and monoterpenes) that dominate biogenic VOCs (BVOCs) have been
60 extensively investigated during the last five decades [e.g., Went, 1960; Rasmussen, 1972;
61 Zimmerman et al., 1979; Lamb et al., 1987; Pierce et al., 1998; Niinemets et al., 1999 and
62 2002; Arneth et al., 2007; Schurgers et al., 2009; Guenther et al., 1995 and 2012]. BVOC
63 emissions were originally computed offline, producing prescribed emission inventories
64 used by regional and global models [e.g., Huang et al., 2011]. However, emissions of
65 BVOCs depend on diurnal, multi-day, and seasonal variations in light intensity,
66 temperature, soil moisture, vegetation type, and leaf area index (LAI) [e.g., Pierce et al.,
67 1998; Niinemets et al., 1999 and 2002; Arneth et al., 2007; Schurgers et al., 2009;
68 Guenther et al., 2012]. Therefore, various **BVOC** emission algorithms have been

Chun Zhao 4/29/2016 10:40 AM

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70 proposed that extrapolate limited laboratory and field measurements to prescribed
71 regional and global ecosystems [e.g., Pierce et al., 1998; Niinemets et al., 1999 and 2002;
72 Arneeth et al., 2007; Schurgers et al., 2009; Guenther et al., 1995 and 2012]. The
73 uncertainties in biogenic emission schemes are mainly due to the scarcity of observations
74 of BVOC fluxes and vegetation distributions over regional scales. Inappropriate coupling
75 strategies between biogenic emission and land-surface schemes may also introduce errors
76 in estimating atmospheric BVOCs. For example, some models specify different
77 vegetation distributions for biogenic emissions and land-atmosphere interaction processes
78 as applied in different parts of models.

79 BVOCs play a significant role in affecting the air quality and regional climate
80 over California, where there have been many studies, such as the Carbonaceous Aerosol
81 and Radiative Effects Study (CARES) [Zaveri et al., 2012] and the California Nexus of
82 Air Quality and Climate Experiment (CalNex) [Ryerson et al., 2013], investigating the
83 impacts of BVOCs and their interaction with anthropogenic pollutants. In the past 20
84 years, California's economy has grown rapidly and the population has increased by 33%
85 [Cox et al., 2009]. Although California has reduced the emissions of most primary
86 pollutants, poor air quality still affects the well-being of millions of people. Nearly all
87 Californians live in areas that are designated as nonattainment for the state (about 99%)
88 and national (about 93%) health-based O₃ and/or PM standards. Accurate predictions of
89 O₃ and PM concentrations are needed to develop effective attainment strategies, but this
90 is complicated, in part, due to uncertainties associated with long-range transport of
91 pollutants and local natural emission sources such as BVOCs.

92 In California, the complex topography and distribution of vegetation makes it

Chun Zhao 4/29/2016 10:40 AM
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94 difficult for models to capture the variability of BVOCs at regional and local scales. For
95 example, Fast et al. [2014] showed that simulated biogenic emissions varied by as much
96 as a factor of 2 within 8 km of an observation site in Cool, California. They also found
97 that daytime mixing ratios of isoprene and monoterpenes from a regional simulation
98 using the Weather Research and Forecasting model with chemistry (WRF-Chem) [Grell
99 et al., 2005; Fast et al., 2006] are usually a factor of two smaller than the observations
100 collected both at the rural Cool site and an urban Sacramento site. Conversely, simulated
101 monoterpene mixing ratios were similar to observations during the day but by a factor of
102 three too high at night at the observation site in Cool. They suggested that the biogenic
103 emission rates calculated based on the Model of Emissions of Gases and Aerosols from
104 Nature version 2.0 (MEGAN v2.0) might contribute to major biases in their simulations.
105 Knote et al. [2014] also found that their simulations using WRF-Chem with MEGAN
106 v2.0 produced BVOC concentrations that were too small over Los Angeles, and
107 suggested that there might be deficiencies in the description of vegetation in urban areas.
108 Thus, it is evident that uncertainties in simulated atmospheric BVOCs can arise from how
109 well vegetation is represented in models. Furthermore, to our knowledge, none of the
110 numerous chemical transport modeling studies for California have investigated the
111 sensitivity of BVOC simulations to land surface schemes and vegetation distributions.
112 To better understand the uncertainties in simulating BVOCs associated with land
113 surface schemes and vegetation distributions in California, the latest version of MEGAN
114 (MEGAN v2.1) is coupled into the CLM4 land surface scheme of WRF-Chem in this
115 study. Multiple sensitivity experiments are conducted using this improved modeling
116 framework at a relatively high spatial resolution to capture the region's complex

Chun Zhao 4/29/2016 10:40 AM

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118 topography and vegetation distribution. Simulations are conducted using WRF-Chem
119 with “fully” coupled version of CLM4 and MEGAN v2.1 (i.e., CLM4 and MEGAN share
120 a consistent vegetation dataset) and compared with the measurements collected during
121 CARES and CalNex conducted in June 2010. This new coupling also adds the capability
122 of quantifying the impact of different vegetation distributions on simulating BVOCs.
123 Simulations are also performed using two land surface [schemes](#) (Noah and CLM4)
124 coupled with MEGAN v2.0. As with previous studies using WRF-Chem, MEGAN v2.0
125 uses a different vegetation dataset from the land surface [schemes](#). The WRF-Chem
126 experiments with MEGAN v2.0 and MEGAN v2.1 are included together here as a
127 reference for future studies in the community and for users interested in migrating from
128 the widely used v2.0 to v2.1.

129 The rest of manuscript is organized as follows. Sections 2 and 3 describe the
130 WRF-Chem model and the observations used in this study, respectively. The sensitivity
131 of modeling BVOCs to the land surface schemes and the vegetation distributions are
132 analyzed in section 4. The findings are then summarized and discussed in section 5.

133

134 **2. Model Description and Experimental Design**

135 **2.1 WRF-Chem**

136 The WRF-Chem (v3.5.1) configuration is similar to that used by Fast et al. [2014]
137 for studying aerosol evolution over California, except that this study excludes aerosols
138 and focuses on simulated BVOCs. The model includes numerous options for the
139 treatment of physics and chemistry processes. In this study, the SAPRC-99
140 photochemical mechanism [Carter, 2000a,b] is selected to simulate gas-phase chemistry,

141 and the Fast-J parameterization [Wild et al., 2000] for photolysis rates. For all the
142 simulations in this study, we use the Yonsei University (YSU) parameterization [Hong et
143 al., 2006] for the planetary boundary layer (PBL), the Monin-Obukhov similarity theory
144 [Paulson, 1970] to represent the surface layer, the Morrison two-moment
145 parameterization [Morrison et al., 2009] for cloud microphysics, the Kain-Fritsch
146 parameterization [Kain 2004] for sub-grid scale clouds and precipitation, the rapid
147 radiative transfer parameterization (RRTMG) for longwave and shortwave radiation
148 [Iacono et al., 2008]. Since Fast et al. [2014] has evaluated the simulated meteorological
149 fields and gases and aerosols with a similar model configuration, this study will focus
150 primarily on the BVOC simulation.

Chun Zhao 4/29/2016 10:40 AM

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151 **2.2 Land surface schemes**

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152 Two land surface schemes, Noah and CLM4.0, are used to quantify how
153 differences in the treatment of land surface processes, including latent and sensible heat
154 fluxes, soil moisture, and surface albedo, affect near-surface meteorological conditions
155 and consequently simulated BVOC emissions and concentrations. The Noah land surface
156 scheme, described by Barlage et al. [2010] and LeMone et al. [2010a, 2010b], has been
157 used in numerous studies with WRF-Chem. Noah has four soil layers, with a total depth
158 of two meters and a single slab snow layer that is lumped with the top-soil layer, which is
159 set to a combined depth of 10 cm. It uses the 24 United States Geological Survey (USGS)
160 land-use types, and does not treat sub-grid scale variability within a model grid cell.

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161 The CLM4 (Community Land Model version 4.0) [Lawrence et al. 2011; Jin et
162 al., 2012] was recently coupled and released with WRF (since v3.5) as one of the land
163 surface scheme options. CLM4 in global and region applications has been shown to be

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168 accurate in describing snow, soil, and vegetation processes [Zeng et al., 2002; Jin and
169 Miller, 2007; Zhao et al., 2014]. CLM4 includes five layers for snow, 10 layers for soil,
170 and a single-layer for vegetation. The soil is divided into 19 categories defined according
171 to percentages of sand and clay. The two-stream approximation [Dickinson, 1983] is
172 applied to vegetation when calculating solar radiation reflected and absorbed by the
173 canopy as well as radiation transfer within the canopy. Each model grid cell can be
174 divided into a maximum of 10 smaller cells to account for sub-grid scale heterogeneity
175 and its impact on the land surface processes. The 24 USGS land-use types are mapped to
176 the 16 plant functional types (PFTs) in CLM4 based on a lookup table derived from
177 Bonan et al. [1996]. Additional technical details of CLM4 are provided in Oleson et al.
178 [2004].

179 2.3 MEGAN and coupling with CLM4

180 MEGAN is a modeling framework for estimating fluxes of biogenic compounds
181 between terrestrial ecosystems and the atmosphere using simple mechanistic algorithms
182 to account for the major known processes controlling biogenic emissions [Guenther et al.,
183 2006, 2012]. Two versions (v2.0 and v2.1) of MEGAN are used in this study. MEGAN
184 v2.1 is an update from MEGAN v2.0 [Guenther et al., 2006; Sakulyanontvittaya et al.,
185 2008] that includes additional compounds, emission types, and controlling processes.
186 MEGAN v2.1 estimates emissions (F_i) for 19 compound classes (i) from terrestrial
187 landscapes based on emission factors ($\epsilon_{i,j}$) at standard conditions for vegetation type j
188 with fractional grid box areal coverage χ_j , i.e., $F_i = \gamma_i \sum \epsilon_{i,j} \chi_j$, where γ_i is emission activity
189 factor from the processes controlling emission responses to environmental and
190 phenological conditions [Guenther et al., 2006, 2012].

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195 For emission factors, MEGAN v2.0 enabled users to customize vegetation
196 emission type schemes ranging from detailed (e.g. individual plant species or sub species)
197 to generic (e.g. a few broad vegetation categories). MEGAN2.1 emission factors can be
198 specified from gridded maps based on species composition and species-specific emission
199 factors or by using PFT distributions and the PFT specific emission factors. MEGAN2.0
200 defines emission factors as the net flux of a compound into the atmosphere, while
201 MEGAN2.1 emission factor represents the net primary emission that escapes into the
202 atmosphere but is not the net flux because it does not include the downward flux of
203 chemicals from above canopy. The difference in the definition (net flux versus primary
204 emission) of emission factors affects the emission factors of compounds with
205 bidirectional exchange but does not impact MEGAN isoprene and monoterpene emission
206 factors because they have small deposition rates relative to emission rates. In this study,
207 both MEGAN v2.0 and v2.1 estimate biogenic species emissions based on the PFT
208 distributions and the PFT specific emission factors. MEGAN v2.0 and v2.1 use 4 and 16
209 PFTs, respectively, as described below in Section 2.4.

210 The publically available version of WRF-Chem includes the MEGAN v2.0
211 scheme for calculating BVOC emission fluxes (WRF-Chem user guide:
212 [http://ruc.noaa.gov/wrf/WG11/Users guide.pdf](http://ruc.noaa.gov/wrf/WG11/Users%20guide.pdf)). It has been widely used for gas and
213 aerosol simulations [e.g., Shrivastava et al., 2011, 2013; Gao et al., 2011, 2014; Knote et
214 al., 2014; Fast et al., 2014]. In the released version, MEGAN v2.0 can be used with any
215 land surface scheme available in WRF-Chem including Noah and CLM4. However,
216 MEGAN v2.0 was originally not coupled into the land surface scheme in WRF-Chem
217 (since v3.1). The biogenic emission calculation in MEGAN uses both instantaneous and

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Deleted:), and used the climatological monthly mean surface air temperature and solar radiation to represent the values for last few days for

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225 ~~the past-days' surface air temperature and solar radiation. MEGAN v2.0 obtains the~~
226 ~~instantaneous value from the land surface scheme and the past-days' value from the~~
227 ~~climatological monthly mean dataset. In contrast, MEGAN v2.1 obtains both values~~
228 ~~directly from CLM.~~ Figure 1 shows the example of the comparison between the input
229 climatological and model simulated monthly mean surface air temperature in June. It is
230 apparent that the monthly-averaged simulated surface air temperature is much different
231 from the climatology value. In addition, the vegetation dataset (referred to as VEG-M,
232 will be discussed in Section 2.4) used in MEGAN v2.0 for calculating BVOC emission
233 fluxes is also different from the one used by the land surface scheme, which allows
234 MEGAN v2.0 to be used with any of the available land surface schemes (e.g., Noah and
235 CLM4) in WRF-Chem. This inconsistency in vegetation distributions may introduce
236 errors in simulating emissions and concentrations of BVOC. To avoid this inconsistency,
237 we have coupled MEGAN v2.1 with WRF-Chem embedded in the CLM4 land surface
238 scheme. Therefore, the coupling of MEGAN v2.1 and CLM4 in WRF-Chem now has the
239 same functionality as CLM4 in the Community Earth System Model (CESM) [Lawrence
240 et al. 2011]. With this coupling strategy, MEGAN v2.1 also uses the same vegetation
241 dataset (i.e., 16 PFTs converted from the USGS dataset as discussed in Section 2.2) that
242 CLM4 uses for all other land surface processes; this means, however, that MEGAN v2.1
243 can only be used with CLM4 in WRF-Chem. In addition, MEGAN v2.1 can compute
244 BVOC emissions that account for the sub-grid variability of vegetation distributions
245 within CLM4.

246 2.4 Vegetation datasets

247 As mentioned previously, the first 16-PFT dataset (referred to as USGS hereafter)

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252 used by CLM4 is converted from the default 24 USGS land cover dataset used by WRF-
253 Chem based on a lookup table derived from Bonan et al. [1996]. This method is also
254 applied to three other 16-PFTs datasets (referred to as VEG1, VEG2, and VEG3) used by
255 CLM4 in WRF-Chem. The sensitivity of simulating BVOC emissions by CLM4 to these
256 four 16-PFTs datasets is quantified. The VEG1, VEG2, and VEG3 datasets are derived
257 from different sources as described next.

258 The VEG1 dataset is from the PFT fractional cover product by Ke et al. [2012],
259 which was developed from the Moderate Resolution Imaging Spectroradiometer
260 (MODIS) PFT classifications for the year 2005 for determining seven PFTs including
261 needleleaf evergreen trees, needleleaf deciduous trees, broadleaf evergreen trees,
262 broadleaf deciduous trees, shrub, grass and crop for each 500 m pixel. The WorldClim 5
263 arc-minute (0.0833°) [Hijmans et al., 2005] climatological global monthly surface air
264 temperature and precipitation data was interpolated to a 500 m grid and used to further
265 reclassify the PFTs into 15 PFTs, and fractions of crop grasses were mapped based on the
266 method presented in Still et al. [2003]. Pixels with barren land and urban areas were
267 reassigned to the bare soil class. The bare soil and the 15 PFTs from the 500-m grid were
268 then aggregated to a 0.05° grid.

269 The VEG2 dataset is obtained from the NCAR CESM data repository [Oleson et
270 al., 2010], available on a 0.05° grid and derived using a combination of the 2001 MODIS
271 Vegetation Continuous Field (VCF), MODIS land cover product for year 2000
272 [Lawrence and Chase, 2006; Lawrence and Chase, 2007], and 1992-1993 AVHRR
273 Continuous Field Tree Cover Project data [Lawrence and Chase, 2007; Lawrence et al.,
274 2011]. The monthly surface air temperature and precipitation data from Willmott and

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277 Matsuura [2001] was used to further reclassify the seven PFTs into bare soil and 15 PFTs
278 in the tropical, temperate and boreal climate groups based on climate rules described by
279 Bonan et al. [2002]. Fractions of crop grasses were mapped based on the method
280 presented in Still et al. [2003].

281 The VEG3 dataset is derived from a high-resolution (30 arc-second) dataset over
282 the U.S. with 16 PFT classifications for the year 2008. The dataset was created by
283 combining the National Land Cover Dataset (NLCD, Homer et al., 2004) and the
284 Cropland Data Layer (see <http://nassgeodata.gmu.edu/CropScape/>), both of which were
285 based on the 30-m LANDSAT-TM satellite data. Vegetation species composition
286 information was obtained from the Forest Inventory and Analysis (see
287 <http://www.fia.fs.fed.us>) and the soil data from the Natural Resources Conservation
288 Services (see <http://sdmdataaccess.nrcs.usda.gov/>). The processing included adjusting
289 the NLCD tree cover estimates in urban areas to account for the substantial
290 underestimation of trees in the LANDSAT-TM data [Duhl et al., 2012]. This was
291 accomplished using the regionally specific adjustment factors for urban NLCD developed
292 by Greenfield et al. [2009] using the high-resolution imagery.

293 Figure 2 shows the spatial distributions of the dominant PFT in each 4×4 km² grid
294 cell of the simulation domain from each of the four datasets. Not only are the grid-
295 dominant PFTs very different among the four datasets, but the sub-grid distributions of
296 PFTs are different as well (not shown). The domain-averaged fractions of 16 PFTs from
297 the four datasets listed in Table 1 also illustrate the differences in PFT distributions. For
298 example, the fraction of temperate broadleaf deciduous tree ranges from 0.4% in VEG1
299 to 1.8% in VEG2 and the fraction of temperate broadleaf deciduous shrub ranges from

300 10.8% in VEG3 to 37.5% in VEG1. In MEGAN v2.0 of WRF-Chem, only four PFTs
301 (refer to VEG-M) that are broadleaf tree, needleleaf tree, shrub, and herbaceous
302 vegetation categories, are considered for the biogenic emission calculation because they
303 are the only ones included in the MEGAN v2.0 PFT scheme. As discussed previously,
304 these are different from the USGS vegetation distribution used by Noah and CLM4 and
305 may cause additional biases. The distributions of the four PFTs used by MEGAN v2.0 are
306 shown in Figure 3. This difference in PFT distributions can affect the BVOC emission
307 calculations, primarily through determining distributions of PFT specific emission factors
308 and leaf area indices (LAI) that are prescribed with PFTs in this study. For example,
309 Figure 4 shows the biogenic isoprene emission factor for each PFT prescribed in
310 MEGAN v2.0 and MEGAN v2.1 in CLM4. In MEGAN v2.1, it shows that temperate
311 broadleaf deciduous tree (PFT 7 listed in Table 1) has a large isoprene emission factor,
312 while temperate needleleaf evergreen tree (PFT 1 listed in Table 1) has a small isoprene
313 emission factor. A similar difference between broadleaf trees and needleleaf trees is
314 indicated for MEGAN v2.0. Figure 5 shows the spatial distributions of averaged biogenic
315 isoprene emission factor used in MEGAN v2.0 and v2.1 with different PFTs. It is evident
316 that the difference in the distributions of PFTs results in a significant difference in spatial
317 distributions of the isoprene emission factor. Figure 6 shows the spatial distributions of
318 LAI used in MEGAN v2.0 and v2.1. The differences in the spatial distributions of LAI
319 can significantly affect the biogenic emission calculation in MEGAN. It should be noted
320 that in MEGAN v2.0 used in WRF-Chem, the LAI used for the calculation of the
321 biogenic emissions is prescribed using the 4 PFTs, which is different than the land
322 scheme that uses the LAI derived from the 24 USGS land categories.

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vegetation categories used by MEGAN v2.0
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v2.0 of WRF-Chem, only these four PFT's are
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calculation because they are the only ones
included in the MEGAN v2.0 PFT scheme.

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these are different from the USGS vegetation
distribution used by Noah and CLM4 and may
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340

341 2.5 Numerical experiments

342 The simulations are performed using a domain encompassing California (Fig. 1)
343 with a horizontal grid spacing of 4 km and 279×279 grid cells (113°W-128°W, 32°N-
344 43°N) and 51 vertical layers up to 100 hPa with about 35 layers below 2 km. The
345 simulation period is from May 25 to June 30 2010, but only the results in June are used
346 for analysis to allow for the model to “spin-up” realistic distributions of trace gases. The
347 initial and boundary conditions are prescribed by large-scale meteorological fields
348 obtained from the North American Regional Reanalysis (NARR) data with updates
349 provided at 6-h intervals, which also provide the prescribed sea surface temperature
350 (SST) for the simulations. The modeled u and v wind components and temperature in the
351 free atmosphere above the planetary boundary layer are nudged towards the NARR
352 reanalysis data with a time scale of 6 hours [Stauffer and Seaman, 1990]. Chemical
353 lateral boundary conditions are from the default profiles in WRF-Chem, which are based
354 on the averages of mid-latitude aircraft profiles from several field studies over the eastern
355 Pacific Ocean [McKeen et al., 2002].

356 Anthropogenic emissions were obtained from the CARB 2008 ARCTAS emission
357 inventory developed for the NASA Arctic Research of the Composition of the
358 Troposphere from Aircraft and Satellite (ARCTAS) mission over California [Pfister et
359 al., 2011]. The CARB inventory contains hourly emissions for a 13-day period using a 4-
360 km grid spacing over California. We created diurnally averaged emissions from 5 of the
361 weekdays and 2 of the weekend days and used those averages for all weekdays and
362 weekends and applied these over the entire simulation period. Anthropogenic emissions

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367 from the 2005 National Emissions Inventory (NEI) (WRF-Chem user guide from
368 http://ruc.noaa.gov/wrf/WG11/Users_guide.pdf) were used for regions outside of
369 California. Biomass burning is not considered in the present study, because satellite
370 detection methods indicated that there were very few fires in California during the
371 simulation period. Biogenic emissions were computed on-line using the MEGAN model
372 and lumped into isoprene, terpenes, and sesquiterpenes for the SAPRC-99 photochemical
373 mechanism.

374 As discussed previously, multiple numerical experiments summarized in Table 2
375 are conducted with different combinations of land surface schemes and vegetation
376 datasets to investigate the sensitivity of BVOC simulation to land surface schemes and
377 vegetation distributions. First, we conduct two experiments using MEGAN v2.0 coupled
378 with the Noah (Mv20Noah) and CLM4 (Mv20CLM) land surface schemes, respectively.
379 The Noah land surface scheme is only coupled with MEGAN v2.0 in WRF-Chem. In
380 these two experiments, the two land surface schemes use the USGS vegetation
381 distributions while MEGAN v2.0 uses a separate vegetation map (VEG-M) to estimate
382 BVOC emissions. By comparing these two experiments, the impact of land surface
383 schemes on simulated BVOC concentrations are examined. Second, we conduct four
384 experiments using MEGAN v2.1 embedded in the CLM4 land surface scheme with four
385 different vegetation datasets, i.e., USGS (Mv21USGS), VEG1 (Mv21V1), VEG2
386 (Mv21V2), and VEG3 (Mv21V3). The differences among these four experiments show
387 the impact of vegetation distributions on simulated BVOC concentrations.

388 We note that MEGAN v2.0 and v2.1 use different vegetation datasets and are
389 implemented in WRF-Chem in different ways, but the objective of this study is not to

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392 explore how the formulations of these two versions of MEGAN affect BVOC
393 concentrations. The better way for exploring the version difference of MEGAN is to
394 implement both versions in the same way and use the same vegetation dataset. The
395 simulated BVOC emissions and concentrations by WRF-Chem with MEGAN v2.0 and
396 MEGAN v2.1 are included together here as a reference for future studies in the
397 community and for users interested in migrating from the widely used v2.0 to v2.1.

398

399 **3. Observations**

400 Measurements of VOCs collected by proton transfer reaction mass spectrometer
401 (PTR-MS) instruments [Lindinger et al., 1998] and a gas chromatography instrument
402 [Gentner et al., 2012] over California during June of 2010 as part of the CARES and
403 CalNex campaigns are used to evaluate the simulated isoprene and monoterpene
404 concentrations. CARES was designed to address science issues associated with the
405 interactions of biogenic and anthropogenic precursors on SOA, black carbon mixing state,
406 and the effects of organic species and aerosol mixing state on optical properties and the
407 activation of cloud condensation nuclei [Zaveri et al., 2012]. As shown in Figure 7,
408 ground-based instruments were deployed at two sites (T0 and T1) in northern California:
409 T0 in Sacramento (38.649 °N, -121.349°W, ~ 30 m m.s.l., denoted by red upward
410 triangle) and T1 in Cool (38.889°N, -120.974°W, ~ 450 m m.s.l., denoted by red
411 downward triangle), a small town located about 40 km northeast of Sacramento. The U.S.
412 Department of Energy (DOE) Gulfstream 1 (G-1) research aircraft sampled
413 meteorological, trace gas, and aerosol quantities aloft in the vicinity of the T0 and T1
414 sites, denoted by black lines in Figure 8. Zaveri et al. [2012] described the

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417 instrumentation for each of the surface sites and Shilling et al. [2013] described VOC
418 measurements on the G-1. Most of the sampling during CARES occurred between 2 and
419 28 June, and only the aircraft sampling within 1 km of the surface is used to evaluate
420 model simulations because G-1 sampled below 1 km for the majority of time.

421 CalNex was designed to address science issues relevant to emission inventories,
422 dispersion of trace gases and aerosols, atmospheric chemistry, and the interactions of
423 aerosols, clouds, and radiation [Ryerson et al., 2013]. Ground-based instruments were
424 deployed at two sites in southern California as shown in Figure 7: one in Pasadena
425 (34.141°N, -118.112°W, ~240 m m.s.l., denoted by the red circle) and one in Bakersfield
426 (35.346°N, -118.965°W, ~ 123 m m.s.l., denoted by the red square). The NOAA WP-3D
427 research aircraft sampled meteorological, trace gas, and aerosol quantities aloft along
428 flight paths shown in Figure 7 (denoted by blue lines). While most of the CalNex aircraft
429 tracks below an altitude of 1 km were conducted in southern California in the vicinity of
430 the Los Angeles basin, the WP-3D also flew within the Central Valley and in the vicinity
431 of Sacramento on some days. A detailed description of the instrumentation for each of the
432 CalNex surface sites and mobile platforms is given by Ryerson et al. [2013]. Most of the
433 sampling during CalNex was conducted before June 16 and only the aircraft sampling
434 below 1 km is used to evaluate the model simulations.

435

436 4. Results

437 4.1 Impact of land surface schemes

438 4.1.1 Biogenic isoprene and monoterpene emissions

439 Figure 7 shows the spatial distributions of biogenic isoprene emissions averaged

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443 over June for the six simulations listed in Table 2. Biogenic isoprene emissions occur in
444 vegetated regions of California with the highest emission rates along the foothills of the
445 Sierra Nevada where oak trees are the dominant plant species. To show the difference in
446 biogenic isoprene emissions among the cases more clearly, Figure 8a and 8b zoom in on
447 the CARES (northern California) and CalNex (southern California) sampling regions,
448 respectively. In both regions the differences in land surface schemes had a relatively
449 small impact on the biogenic isoprene emissions over California in terms of both spatial
450 distribution and magnitude, although the emissions from Mv20CLM were a little larger
451 than those from Mv20Noah. The domain summed biogenic isoprene emissions for the
452 entire month of June from Mv20Noah and Mv20CLM are 1.4×10^9 and 1.6×10^9 mole,
453 respectively. Figure 9a and 9b are similar to Figure 8a and 8b, except that biogenic
454 monoterpene emission fluxes are shown. In general, the spatial patterns of emissions of
455 the two biogenic species are similar, except that the peak areas of monoterpene emissions
456 are shifted slightly. For example, the peak monoterpene emissions in northern California
457 occur further northeast at higher elevations of the Sierra Nevada that are dominated by
458 needleleaf evergreen trees. The impact of land surface schemes on biogenic monoterpene
459 emissions is also small over California in terms of both spatial patterns and magnitudes,
460 although the emissions from Mv20CLM are a little larger than those from Mv20Noah.
461 The domain summed biogenic monoterpene emissions for the entire month of June from
462 Mv20Noah and Mv20CLM are 1.0×10^8 and 1.1×10^8 mole, respectively.

463 The similarity in estimating biogenic emissions between the experiments with two
464 land surface schemes is also summarized in Figures 10 and 11, which show the average
465 diurnal biogenic isoprene and monoterpene emission rates at the four observation sites.

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475 The similarity between Mv20Noah and Mv20CLM (red and orange lines) is likely due to
476 the same vegetation map in MEGAN v2.0 to estimate biogenic emissions. Although the
477 two land surface schemes produce slightly different values of surface temperature (Fig.
478 1), soil moisture (not shown), and net solar radiation near the surface (not shown), their
479 impact on the biogenic emissions was small. Both BVOC species have peak emission
480 rates in the early afternoon. One noteworthy difference in diurnal variation of the two
481 biogenic species emission rates is that there is no isoprene emitted during the night while
482 the amount of monoterpenes emitted during the night is small but not negligible. This can
483 contribute to differences in the diurnal variation of the mixing ratios of two biogenic
484 species, as will be discussed next.

485 4.1.2 Isoprene and monoterpene mixing ratios

486 Figures 12a,b and 13a,b show the spatial distributions of monthly-averaged
487 surface mixing ratios of isoprene+MVK(methyl-vinylketone)+MACR(methacrolein) and
488 monoterpenes, respectively, around the CARES (northern California) and the CalNex
489 (central and southern California) sampling regions simulated by the six experiments
490 listed in Table 2. Due to the fast chemical transition from isoprene to MVK and MACR,
491 the sum of isoprene+MVK+MACR mixing ratios can better reflect the impact of
492 biogenic isoprene emissions than isoprene mixing ratio alone [Shilling et al., 2013]. In
493 general, the spatial patterns and magnitudes of surface isoprene+MVK+MACR and
494 monoterpene mixing ratios over the two regions are similar from the two MEGAN v2.0
495 experiments with the Noah and CLM4 land surface schemes, respectively. The spatial
496 patterns of surface mixing ratios of isoprene+MVK+MACR and monoterpenes are
497 similar to the spatial variability in the emission rates.

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502 There is difference between the two experiments at specific locations, which is
503 partly reflected in the comparison of average diurnal variations of surface mixing ratios
504 of isoprene+MVK+MACR and monoterpenes at the four observation sites shown in
505 Figure 14 and Figure 15. At the Bakersfield site, only isoprene mixing ratios were
506 reported so that the comparison is for isoprene only. Note that the values for the
507 Bakersfield and Pasadena sites are averaged over the first two weeks of June to be
508 consistent with the observations. Although both experiments with Noah and CLM4 (red
509 and orange lines, respectively) simulate similar isoprene emission fluxes with the
510 maximum in the afternoon (Fig. 10), their respective isoprene+MVK+MACR mixing
511 ratios are different at the four sites, particularly at site T0, where the Mv20CLM
512 simulated isoprene+MVK+MACR mixing ratios during the daytime are about a factor of
513 2 larger than those from Mv20Noah. This inconsistence mainly results from the
514 differences in the near surface meteorology, such as net surface radiation and
515 temperature, between the two experiments (not shown) that affects photochemistry, but
516 this impact of surface meteorology occurs only at limited locations. When compared to
517 the observations, both experiments significantly underestimate the
518 isoprene+MVK+MACR mixing ratios except at the Bakersfield site. Figure 15 is
519 identical to Figure 14, except for surface monoterpene mixing ratios. Note that there were
520 no monoterpene data reported for the Bakersfield and Pasadena sites, so only the
521 simulation results are shown. In contrast to isoprene+MVK+MACR, monoterpenes
522 exhibit peak surface mixing ratios during the nighttime due to the strong photolysis
523 activity that makes the lifetime of monoterpenes short during the daytime and the small
524 emissions into a shallow boundary layer during the nighttime (Fig. 11). In general, the

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535 difference between the Mv20Noah and MV20CLM experiments in monoterpene mixing
536 ratios is relatively small at these four sites, particularly during the daytime. When
537 compared to the observations, both experiments overestimate the diurnal variation and
538 the nighttime surface monoterpene mixing ratios at the T0 and T1 sites.

539 Figures 16 and 17 show the comparison of the observed and simulated mixing
540 ratios of isoprene+MVK+MACR and monoterpenes, respectively, along the G-1 and
541 WP-3D flight tracks below 1 km. Model results are sampled along the flight tracks. As
542 shown in Figure 7, the G-1 flight mainly flew over northern California around the T0 and
543 T1 sites, while the WP-3D flew over a larger area covering both southern California and
544 the Central Valley. To better reflect the spatial variability in the BVOCs, the flight tracks
545 of both flights are separated into two regions as indicated by the black lines in Figure
546 12a,b and Figure 13a,b. For the G-1, the flight paths are divided into regions of southwest
547 and northeast of the black line shown in Figures 12a and 13a that is parallel to the Sierra
548 Nevada. The two regions have significantly different vegetation (Fig. 2) resulting in large
549 differences in biogenic emissions. For the WP-3D, the flight paths are divided into
550 regions of south and north of the black line shown in Figures 12b and 13b to separate
551 southern California and the Central Valley. Over southern California, the measured
552 isoprene+MVK+MACR mixing ratios by the PTR-MS over the WP-3D are the upper
553 limit since the PTR-MS may have a small interference in urban areas for isoprene and
554 MVK+MACR.

555 In Figure 16, it is interesting to note that both experiments Mv20Noah and
556 MV20CLM reasonably capture the variability seen in the G-1 isoprene+MVK+MACR
557 measurements over the southwest region even though they underestimate the surface

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569 observations by as much as a factor of 2 at the T0 site (Fig. 14). While both experiment
570 mixing ratios are slightly smaller than observed, the Mv20CLM simulated mixing ratios
571 are a little larger than those from Mv20Noah and closer to the observations. Over the
572 northeast region, both experiments produced similar mixing ratios that were significantly
573 smaller than the observations, which is consistent with the comparison between the
574 simulated and observed isoprene+MVK+MACR at the T1 site (Fig. 14). As shown in
575 Figure 16, the Mv20CLM simulation produced somewhat larger isoprene+MVK+MACR
576 mixing ratios than Mv20Noah in both southern California and the Central Valley. This is
577 consistent with the comparison at the Bakersfield and Pasadena surface sites. Both
578 simulations also underestimate and overestimate the isoprene+MVK+MACR mixing
579 ratios over southern California and the Central Valley, respectively. The comparison of
580 isoprene+MVK+MACR with aircraft observations may suggest that both experiments
581 underestimate biogenic isoprene emissions over the forested foothills of Sierra Nevada
582 and southern California around Los Angeles, but overestimate the emissions over the
583 Central Valley. The model biases may also be affected, to some extent, by anthropogenic
584 emissions with large uncertainties and the associated non-linear chemistry due to the
585 mixing of anthropogenic and biogenic plumes [Fast et al., 2014].

586 Figure 17 shows that both experiments Mv20Noah and Mv20CLM significantly
587 underestimate the monoterpene mixing ratios over all the regions sampled by the G-1 and
588 WP-3D aircraft and that the differences between the simulations were negligible. The
589 average monoterpene mixing ratios sampled by the G-1 below 1 km was comparable to
590 the surface measurement at the T0 site during the daytime, but somewhat higher than the
591 observations at the T1 site. The simulated mixing ratios averaged along the flight tracks

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597 were much smaller than those at the two surface sites, suggesting that it may be difficult
598 for model to simulate the large spatial heterogeneity of the monoterpene mixing ratios.
599 This could result from the biases in biogenic monoterpene emissions and/or the chemical
600 mechanism for monoterpene oxidation and how chemistry is coupled with turbulent
601 mixing within the simulated convective boundary layer. It also needs to be noted that the
602 G-1 and WP-3D measured monoterpene mixing ratios are generally below the Limit Of
603 Detection (LOD) of instruments (0.1-0.3 ppbv). Therefore, the true monoterpene mixing
604 ratios could be range between 0 ~ 0.1-0.3 ppbv, which may also contribute to the
605 discrepancy between observations and simulations.

606

607 4.2 Impact of vegetation distributions

608 4.2.1 Biogenic isoprene and monoterpene emissions

609 | Figures [8a,b](#) and [9a,b](#) show that the differences in biogenic isoprene and
610 | monoterpene emission distributions due to using the various vegetation datasets are larger
611 | than the differences resulting from the two land surface [schemes](#). The domain summed
612 | biogenic isoprene emissions for the entire month of June are 2.3, 0.76, 1.7, and 0.92
613 | ($\times 10^9$ mole) from the experiments of Mv21USGS, Mv21V1, Mv21V2, and Mv21V3,
614 | respectively, and biogenic monoterpene emissions are 2.5, 1.7, 1.9, and 1.1 ($\times 10^8$ mole)
615 | from the four experiments, respectively. Each of the four simulations produces high
616 | biogenic isoprene and monoterpene emission rates along the Sierra Nevada that is
617 | covered mainly by oak and pine forests. However, the different forest classifications and
618 | their coverage (Table 1) produce different biogenic isoprene and monoterpene emission
619 | rates along the Sierra Nevada. Another distinct difference among these four simulations

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622 is found over the Central Valley, where the Mv21V1 and Mv21V3 experiments produce
623 significantly lower biogenic isoprene and monoterpene emissions than the Mv21USGS
624 and Mv21V2 experiments. This results from their different spatial distributions of
625 vegetation types. For example, the vegetation dataset in MV21USGS assigns a relatively
626 larger fraction of vegetation over the Central Valley to broadleaf trees, which are biggest
627 contributors of isoprene emissions (Fig. 4).

628 The differences in the spatial distributions of biogenic isoprene and monoterpene
629 emissions due to various vegetation distributions is also illustrated by the average diurnal
630 biogenic isoprene emission rates at the four observation sites shown in Figures 10 and 11.

631 For example, the Mv21V3 simulation produces the largest biogenic isoprene and
632 monoterpene emissions at three of the sites. At the T1 site over the forested foothills of
633 the Sierra Nevada, the Mv21USGS and Mv21V3 simulations produce much larger
634 biogenic isoprene emissions than Mv21V1 and Mv21V2. Even though forest is the
635 dominant vegetation type along the foothills of the Sierra Nevada in all four vegetation
636 datasets (Fig. 2), their different forest classifications and coverage result in biogenic
637 isoprene emission rates that differ by as much as a factor of 8 at the T1 site. Similar to
638 isoprene emissions, the Mv21USGS simulation produces the largest monoterpene
639 emissions at the T1 site. However, the differences in monoterpene emissions among the
640 four vegetation dataset experiments are smaller overall than that for biogenic isoprene
641 emissions. Different vegetation distributions for a typical urban area can also lead to
642 differences in biogenic isoprene and monoterpene emissions. For example at the urban
643 T0 and Pasadena sites, biogenic isoprene and monoterpene emission rates are almost 0 in
644 the Mv21USGS and Mv21V1 experiments, while the rates were significant larger in the

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648 Mv21V3 experiment. This could have profound implications on local oxidant chemistry
649 influencing urban air quality.

650 4.2.2 Isoprene+MVK+MACR and monoterpene mixing ratios

651 As expected, the differences in biogenic isoprene and monoterpene emissions
652 among the four different vegetation distribution experiments lead to large differences in
653 the simulated surface isoprene+MVK+MACR and monoterpene mixing ratios (Figs.
654 [12a,b](#) and [13a,b](#)). Although all the four experiments simulate highest biogenic
655 isoprene+MVK+MACR and monoterpene mixing ratios along the forested foothills of
656 Sierra Nevada, the Mv21V1 and Mv21V3 experiments have the lowest
657 isoprene+MVK+MACR and monoterpene mixing ratios, respectively, corresponding to
658 their lowest biogenic emission rates. Over the Central Valley, Mv21USGS and Mv21V2
659 experiments produce significantly higher isoprene+MVK+MACR mixing ratios than the
660 other two experiments, while Mv21V3 simulates the lowest monoterpene mixing ratios
661 among all the experiments.

662 At the T1 site located in the forested foothills of Sierra Nevada, the Mv21V1
663 simulation produces the lowest isoprene+MVK+MACR mixing ratios (Fig. [14](#)),
664 significantly underestimating the peak concentrations during the day. In contrast, the
665 Mv21USGS and Mv21V3 simulations reasonably capture the observed
666 isoprene+MVK+MACR mixing ratios during the daytime. All four experiments
667 underestimate the isoprene+MVK+MACR mixing ratios by about a factor of 2 during the
668 night. This may indicate that the transported isoprene+MVK+MACR from the
669 surrounding areas of T1 was too low. The negative biases of simulated
670 isoprene+MVK+MACR mixing ratios over the areas surrounding T1 can be reflected by

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674 | Figure [16](#) that shows all the four experiments significantly underestimate the observed
675 | isoprene+MVK+MACR mixing ratios below 1 km in the northeast area around the T1
676 | site (Fig. [12a](#)). Figure [16](#) also shows that Mv21USGS and MV21V3 simulate larger
677 | isoprene+MVK+MACR mixing ratios averaged over the northeast region of northern
678 | California than Mv21V1 and Mv21V2. All four experiments produce similar surface
679 | monoterpene mixing ratios, which are smaller than that from the Mv20Noah and
680 | Mv20CLM with MEGAN [v2.0](#) and are closer to the observed values particularly during
681 | the night. This is consistent with their much lower biogenic monoterpene emissions
682 | during the night (Fig. [11](#)). The four experiments with MEGAN [v2.1](#) simulate higher
683 | daytime monoterpene mixing ratios averaged along the flight tracks below 1 km than the
684 | two experiments with MEGAN [v2.0](#). The simulated mixing ratios are still much lower
685 | than the aircraft observations, although the simulated surface mixing ratios are higher
686 | than the observations at the T1 site (Fig. [15](#)). However, the aircraft measured
687 | monoterpene mixing ratios may also be higher than the true values due to the LOD of
688 | instruments (0.1-0.3 ppbv).

689 | At the T0 site, an urban site, the vegetation coverage in both the Mv21USGS and
690 | Mv21V1 experiments is small so that the isoprene+MVK+MACR and monoterpene
691 | mixing ratios are significantly lower than observed during the daytime. The Mv21V2 and
692 | Mv21V3 experiments reasonably simulate isoprene+MVK+MACR mixing ratios during
693 | the daytime. Over the area surrounding the T0 site (i.e., the southwest area in Fig. [12a](#)), it
694 | is interesting to note that the Mv21USGS and Mv21V2 simulations produced larger
695 | isoprene+MVK+MACR mixing ratios than Mv21V1 and Mv21V3 and closer to the
696 | observations (Fig. [16](#)). This is mainly due to the relatively large isoprene+MVK+MACR

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707 | mixing ratios over the northwest corner of CARES sampling region (Fig. 12a) in the
708 | Mv21USGS and Mv21V2 simulations, consistent with the distributions of biogenic
709 | isoprene emissions over the region. The Mv21V2 and Mv21V3 simulations produced
710 | higher monoterpene mixing ratios than Mv21USGS and Mv21V1, but are still smaller
711 | than the observed values during the daytime not only for the T0 site but also for the
712 | region surrounding T0 as shown in Figure 17.

713 | At the Bakersfield site, the experiments often simulate significantly larger
714 | isoprene mixing ratios than the observations, except for the Mv21V1 simulation that was
715 | always too small. The Mv21V3 simulation produced the highest isoprene mixing ratios
716 | among the experiments. This is consistent with its biogenic isoprene emission rates (Fig.
717 | 10). In addition, the observed surface isoprene mixing ratios show negligible diurnal
718 | variation in contrast to the experiments that produced larger diurnal variations. The
719 | Mv21V3 simulation produced peak isoprene mixing ratios during the daytime that were
720 | likely controlled by its large daytime local biogenic isoprene emission rates (Fig. 10).
721 | The Mv21USGS and Mv21V2 simulations produced peak isoprene mixing ratios during
722 | the early evening, possibly the result of chemistry and transport from regions with higher
723 | biogenic emissions. All four experiments produce small diurnal variation of surface
724 | monoterpene mixing ratios. The Mv21USGS and Mv21V3 simulations produce larger
725 | monoterpene mixing ratios than the other two, consistent with their local emission rates
726 | (Fig. 11).

727 | At the Pasadena site, the Mv21V3 simulation reproduces the observed diurnal
728 | variation of isoprene+MVK+MACR mixing ratios reasonably well. This is consistent
729 | with the area surrounding the Pasadena site, in which the Mv21V3 simulation produces

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737 the largest mixing ratios of isoprene+MVK+MACR both at the surface (Fig. 12b) and
738 aloft (Fig. 16) in the vicinity of Los Angeles. The other three experiments simulated
739 significantly smaller mixing ratios of isoprene+MVK+MACR. Although the values from
740 the other three experiments are still smaller than the observations, they are much closer to
741 the aircraft measurements (within a factor of 2) than at the Pasadena site (Fig. 14).
742 Among the four vegetation sensitivity simulations, Mv21V3 produces higher surface
743 monoterpene mixing ratios than the other three experiments, consistent with their
744 emission rates (Fig. 11). All four vegetation sensitivity experiments produced much
745 lower monoterpene mixing ratios below 1 km (Fig. 17), compared to the aircraft
746 measurements over southern California that may overestimate the true values due to the
747 LOD of instruments (0.1-0.3 ppbv).

748 As discussed previously, all four experiments simulate significantly different
749 isoprene+MVK+MACR and monoterpene mixing ratios over the Central Valley (Figs.
750 12a,b and 13a,b). The Mv21USGS and Mv21V2 simulations produce much larger
751 isoprene+MVK+MACR mixing ratios (0.6 ppbV and 0.5 ppbV, respectively) over the
752 Central Valley than the observed values (~0.1 ppbV). The Mv21V1 and Mv21V3
753 simulations produce monoterpene mixing ratios much closer to observed values. This
754 may indicate that the fraction of broadleaf trees (the main emitter over the region) over
755 the Central Valley from the vegetation datasets USGS and VEG2 are overestimated or
756 the biogenic emission factors estimated for the broadleaf trees are overestimated for this
757 area. For monoterpenes, the Mv21V3 simulation was much smaller than observed, while
758 the mixing ratios from the other three experiments were more comparable. This suggests
759 that the fraction of vegetation emitting monoterpenes is significantly underestimated over

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767 this area in the VEG3 dataset.

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769 **5. Summary and discussion**

770 In this study, the latest version of MEGAN (v2.1) is coupled within the CLM4
771 land scheme as part of WRF-Chem. Specifically, MEGAN v2.1 is implemented into the
772 CLM4 scheme so that a consistent vegetation map can be used for estimating biogenic
773 VOC emissions as well as surface fluxes. This is unlike the older version of MEGAN
774 (v2.0) in the public-released WRF-Chem that uses a standalone vegetation map that
775 differs from what is used in land surface schemes. With this improved WRF-Chem
776 modeling framework coupled with CLM4-MEGAN v2.1, the sensitivity of biogenic VOC
777 emissions and hence of atmospheric VOC mixing ratios to vegetation distributions is
778 investigated. The WRF-Chem simulations are also conducted with the two land surface
779 schemes, Noah and CLM4, with the MEGAN v2.0 scheme for biogenic emissions in each
780 case. The comparison between the Noah and CLM4 driven MEGAN v2.0 biogenic
781 emissions not only serves for investigating the impact of different land surface schemes
782 on the emissions but also provides a reference for all previous studies that used the Noah
783 land surface scheme. Experiments are conducted for June 2010 over California,
784 compared with the measurements from the CARES and CalNex campaigns. The main
785 findings about the modeling sensitivity to the land surface schemes and vegetation
786 distributions include:

- 787 • The WRF-Chem simulation with the CLM4 land surface scheme and the MEGAN
788 v2.0 module (Mv20CLM) produces similar biogenic isoprene and monoterpene
789 emissions in terms of spatial patterns, magnitudes, and diurnal variations as the one

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791 with the Noah land surface scheme (Mv20Noah) in June over California. The
792 similarity in the biogenic emissions between the experiments using two different land
793 schemes is primarily because of using MEGAN v2.0 and the same vegetation map in
794 the two experiments. The spatial patterns and magnitudes of surface
795 isoprene+MVK+MACR and monoterpene mixing ratios are generally similar
796 between the two experiments with the Noah and CLM4 land surface schemes,
797 although there are significant differences at some specific locations due to their
798 differences in the near surface meteorology such as surface net radiation and
799 temperature. Compared with surface and aircraft measurements, both experiments
800 generally underestimate the daytime mixing ratios of isoprene+MVK+MACR but
801 overestimate the nighttime mixing ratios of monoterpenes.

802 • The experiments with the four vegetation datasets result in much larger differences in
803 biogenic isoprene and monoterpene emissions than the ones with the two land surface
804 schemes. The simulated total biogenic isoprene and monoterpene emissions over
805 California can differ by a factor of 3 among the experiments and the difference can be
806 even larger over specific locations. The comparison of mixing ratios of
807 isoprene+MVK+MACR and monoterpenes with the observations indicates the
808 simulation biases can be largely reduced with accurate vegetation distributions over
809 some regions of California. For example, at an observation site at the forested
810 foothills of Sierra Nevada, two experiments with the vegetation distributions from the
811 USGS and VEG3 datasets capture the observed daytime surface mixing ratios of
812 isoprene+MVK+MACR well, with values that are much larger than the experiments
813 with the other two vegetation datasets.

814 • Although vegetation distributions from some datasets do significantly improve the
815 model performance in simulating BVOC mixing ratios more than others, the optimal
816 vegetation dataset cannot be determined, because the improvement by vegetation
817 datasets has dependence on both the region and BVOC species of interest. For
818 example, over the Central Valley, the experiments with the VEG1 and VEG3
819 vegetation datasets simulate isoprene+MVK+MACR mixing ratios that are much
820 closer to observations than the USGS and VEG2 datasets, while the VEG3 dataset
821 significantly underestimates the observed monoterpene mixing ratios. Large biases
822 over some regions of California in all the experiments with current vegetation
823 datasets imply that more effort is needed to improve land cover datasets and/or
824 biogenic emission factors.

825 There are still some large biases existing over some regions of California
826 regardless of the vegetation distributions. For example, all the experiments significantly
827 underestimate the observed isoprene+MVK+MACR mixing ratios below an altitude of 1
828 km over the forest-covered Sierra Nevada. Over the Pasadena area, all the experiments
829 simulate significantly smaller monoterpene mixing ratios than observed. The biases in
830 BVOCs identified in this study may be partly due to inaccurate vegetation distributions in
831 all the vegetation distribution datasets. The biases can also result from the uncertainties in

832 BVOC emission factors for the individual types of vegetation commonly found in
833 California. The constraints on BVOC emission factors applied in models are limited due
834 to sparse measurements of BVOC emission fluxes. The MEGAN scheme in WRF-Chem
835 uses the global averaged emission factors for BVOC emissions for each PFT. Over
836 California, the broadleaf temperate trees are primarily oaks that have relatively higher

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842 | BVOC emission factors compared to the global mean values for temperate broadleaf
843 trees. In addition, the needleleaf trees are pines that have relatively larger monoterpene
844 emission factors compared to global mean values. These biases in emission factors may
845 partly explain why all the experiments generally underestimate mixing ratios of
846 isoprene+MVK+MACR and monoterpenes over the regions with large amounts of trees.
847 The MEGAN scheme using the location-specified emission factor maps that accounts for
848 species composition of trees may provide a better estimate on regional scales.

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849 This study demonstrates large difference between the experiments with the two
850 versions of MEGAN (v2.0 versus v2.1), and that MEGAN v2.1 results in a better
851 comparison with the observations over some parts of the study domain. However, this
852 difference should not be fully attributed to the improvement of MEGAN between the two
853 versions, because the two versions also use different vegetation distributions. The results
854 highlight the importance of sub-grid vegetation distributions in simulating biogenic
855 emissions even at a relatively high horizontal grid spacing (e.g., 4 km in this study). The
856 biogenic emissions can be significantly different even though the dominant vegetation
857 within a model grid box is similar. The comparison of the simulations and the
858 observations at the surface sites and along the aircraft tracks reflects the large spatial
859 variability of biogenic emissions and BVOC mixing ratios over California. It is
860 challenging for model to capture such a spatial heterogeneity of BVOCs if the vegetation
861 distributions are not appropriately represented in the simulation. The relatively large
862 LOD of instruments on the aircrafts for monoterpenes compared to the true
863 concentrations also make the evaluation of simulated monoterpenes difficult. Over a
864 region with relatively low monoterpene concentrations, an instrument with lower LOD is

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870 needed. It is also noteworthy that this study is in a relatively dry and warm season;
871 therefore the impact of biogenic emission treatments may change for other seasons and
872 during periods with higher cloudiness. A multiple-season investigation may be needed in
873 future. Finally, it is also noteworthy that factors other than biogenic emissions can
874 influence the simulated BVOC mixing ratios over California, such as anthropogenic
875 emissions and the oxidation mechanism of BVOCs used in simulations. Therefore,
876 additional direct measurements of biogenic emission fluxes are needed for a better
877 evaluation of simulated BVOC fluxes.

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879 **Code availability**

880 The WRF-Chem version 3.5.1 release can be obtained at
881 http://www2.mmm.ucar.edu/wrf/users/download/get_source.html. Code modifications for
882 implementing MEGANv2.1 into CLM are available upon request by contacting the
883 corresponding author and will be released to public WRF-Chem version.

884

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Table 1 Average percentage of **PFTs** over the simulation domain

PFT # and description	USGS	VEG1	VEG2	VEG3
0 Bare soil	26.0	7.6	38.1	41.6
1 Needleleaf evergreen tree – temperate	13.0	12.5	9.1	10.7
2 Needleleaf evergreen tree - boreal	0.0	0.1	0.0	4.9
3 Needleleaf deciduous tree – boreal	0.1	0.0	0.0	0.0
4 Broadleaf evergreen tree – tropical	0.0	0.0	0.0	0.0
5 Broadleaf evergreen tree – temperate	0.0	0.4	1.9	0.0
6 Broadleaf deciduous tree – tropical	2.9	0.0	0.0	0.0
7 Broadleaf deciduous tree – temperate	1.5	0.4	1.8	1.5
8 Broadleaf deciduous tree – boreal	0.0	0.0	0.0	0.3
9 Broadleaf evergreen shrub - temperate	21.1	5.3	0.0	0.3
10 Broadleaf deciduous shrub – temperate	20.0	37.5	27.4	10.8
11 Broadleaf deciduous shrub – boreal	0.9	0.2	0.0	1.0
12 C₃ arctic grass	0.0	0.0	1.2	2.2
13 C₃ grass	1.0	28.0	14.9	18.9
14 C₄ grass	10.4	0.0	0.0	0.0
15 Crop	3.2	6.5	4.1	6.3

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¹USGS is the 16-PFT dataset converted from the default 24 USGS land cover dataset based on a lookup table derived from Bonan et al. [1996];
²VEG1 is from the PFT fractional cover product by Ke et al. [2012];
³VEG2 is obtained from the NCAR CESM data repository [Oleson et al., 2010];
⁴VEG3 is derived from a dataset over the U.S. with 16 PFT classifications by combining the National Land Cover Dataset (NLCD, Homer et al., 2004) and the Cropland Data Layer (see <http://nassgeodata.gmu.edu/CropScape/>).

Table 2 Experiments of WRF-Chem

	Surface scheme	BVOC scheme	Plant Function Type Dataset				
			USGS/VEG-M	USGS	VEG1	VEG2	VEG3
WRF-Chem	CLM4.0	MEGANv2.0	Mv20CLM	-	-	-	-
	Noah	MEGANv2.1	-	Mv21USGS	Mv21V1	Mv21V2	Mv21V3
	Noah	MEGANv2.0	Mv20Noah	-	-	-	-

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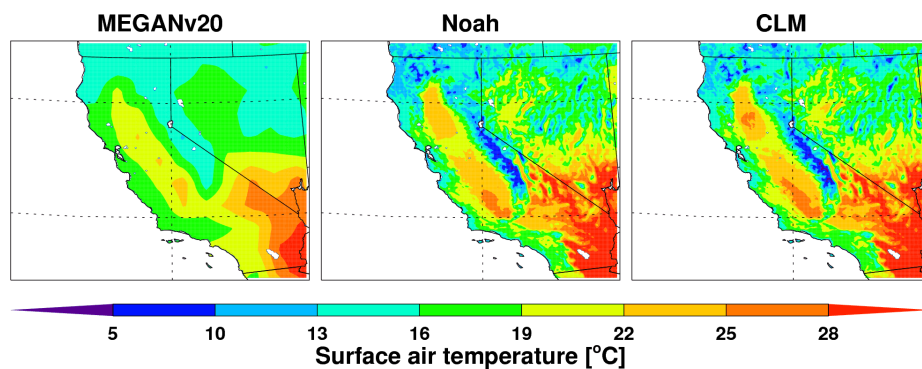
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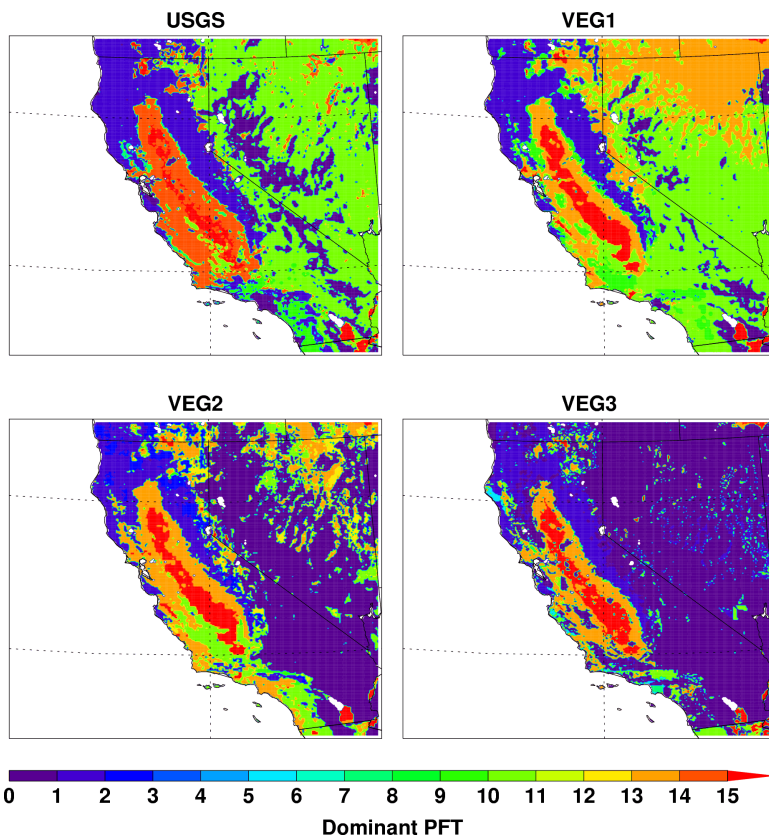


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1226 **Figure 1** Spatial distributions of monthly mean surface air temperature in June 2010 from
1227 the MEGAN v2.0 climatology dataset (MEANv20, prescribed) and the WRF-Chem
1228 simulations with the Noah (Noah, simulated) and CLM4 (CLM, simulated) land surface
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Figure 2. Spatial distribution of dominant PFTs over the simulation domain from the four datasets: USGS, VEG1, VEG2, and VEG3. The PFT number is referred to the list in Table 1.

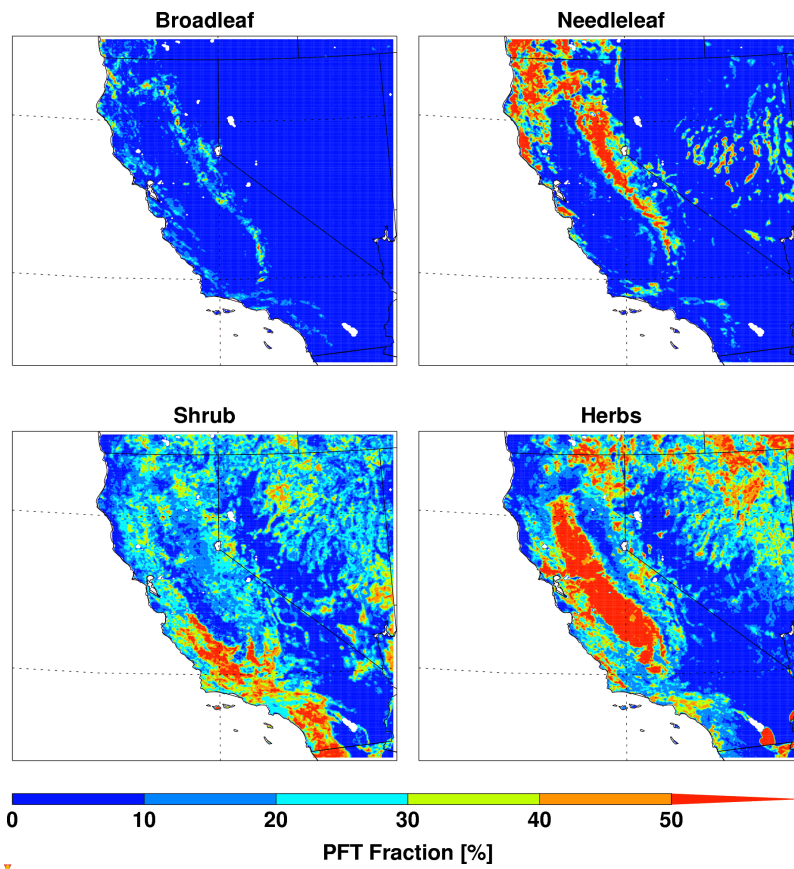
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Figure 3.

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1278 | Spatial distribution of percentage of the four PFTs from the VEG-M used by MEGAN
1279 | v2.0 over the simulation domain.

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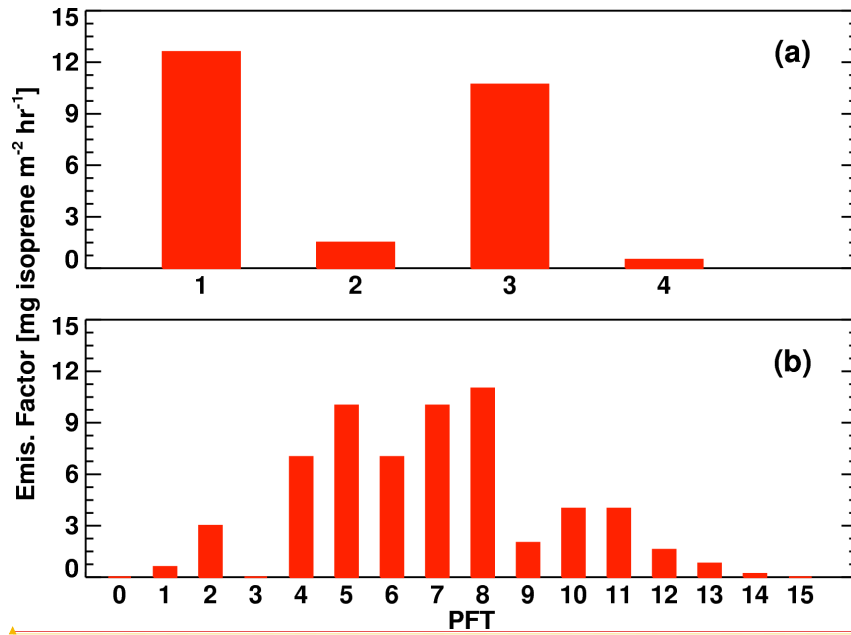
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1292 **Figure 4.** Biogenic isoprene emission factor for each PFT in (a) MEGAN v2.0, the PFT
 1293 number 1-4 is referred to Broadleaf, Needleleaf, Shrub, and Herbs, respectively; (2)
 1294 MEGAN v2.1, the PFT number 0-15 is referred to the list in Table 1.

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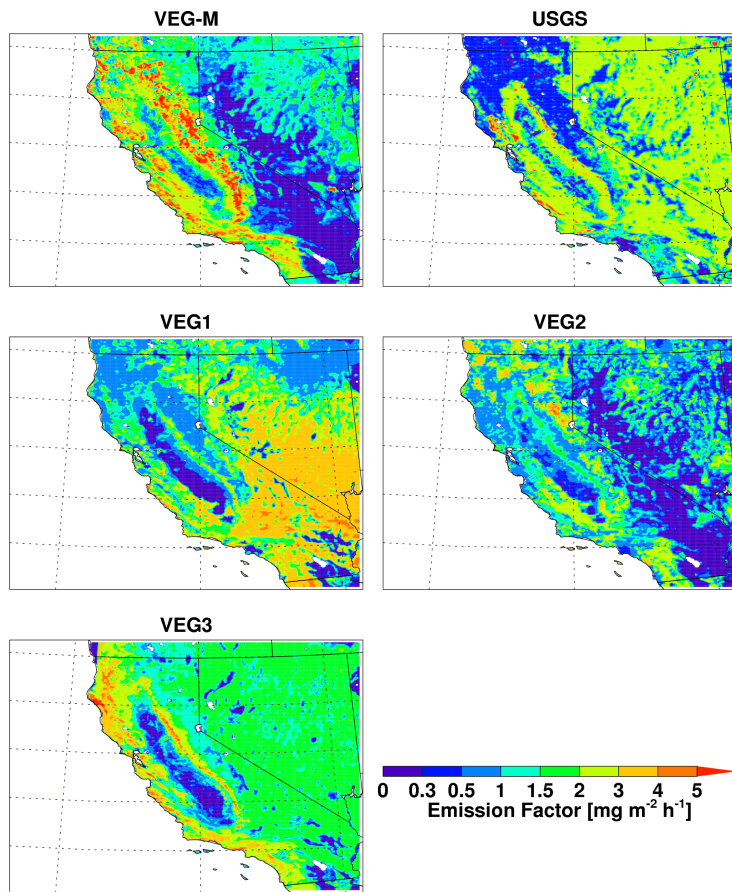
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1305 **Figure 5.** Spatial distribution of PFT-weighted mean biogenic isoprene emission factor
 1306 derived with the VEG-M in MEGAN v2.0 and the USGS, VEG1, VEG2, and VEG3 in
 1307 MEGAN v2.1.
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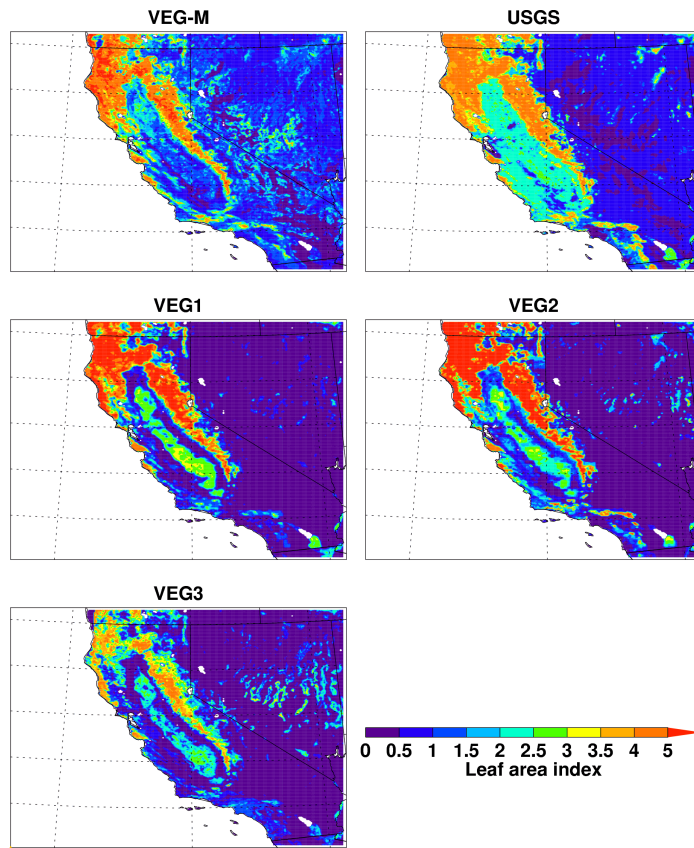
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1318 **Figure 6.** Spatial distribution of leaf area index (LAI) from the VEG-M in MEGAN v2.0
1319 and from the USGS, VEG1, VEG2, and VEG3 in MEGAN v2.1.
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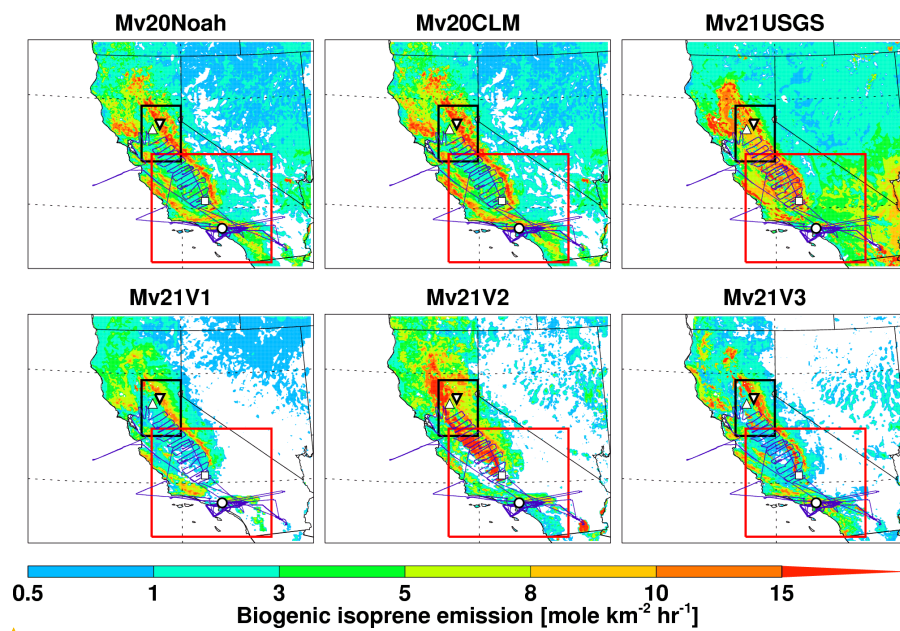
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1330 **Figure 7.** Spatial distributions of biogenic isoprene emissions averaged in June estimated
1331 in the six simulations as listed in Table 2. The four observation sites are shown as T0
1332 (white upward triangle), T1 (white downward triangle), Bakersfield (white square), and
1333 Pasadena (white circle). The CalNex WP-3D flight tracks below 1 km (blue line) during
1334 June 2010 are also shown. The black and red boxes denote the predominant CARES and
1335 CalNex regions, respectively.

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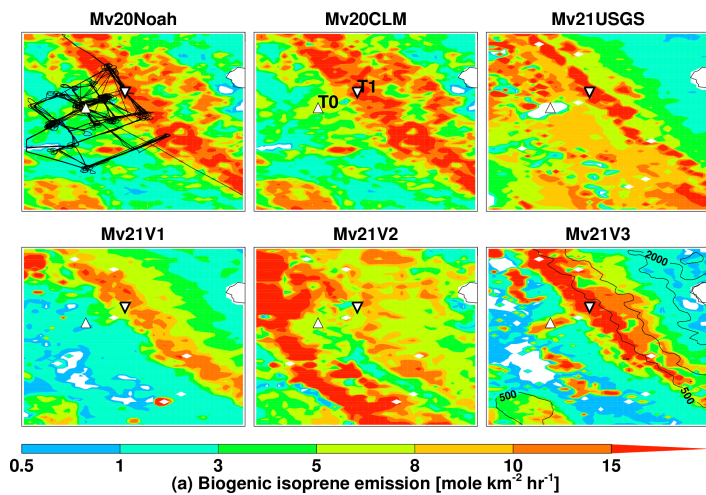
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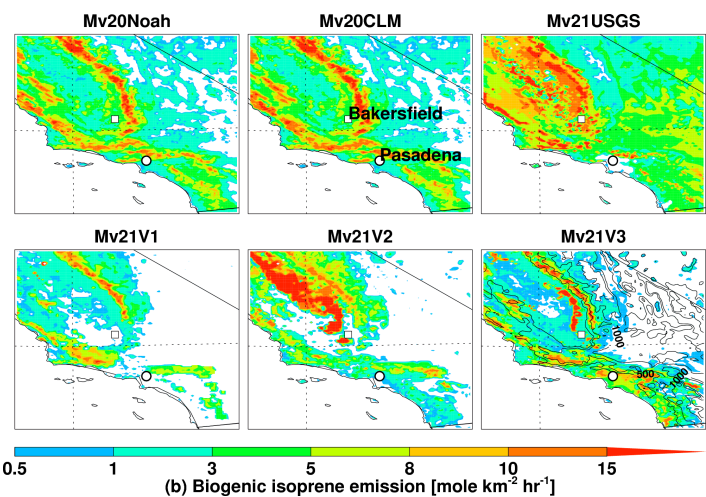
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1346 **Figure 8.** a) Spatial distributions of biogenic isoprene emissions around the CARES
1347 observational sites T0 and T1 (the black box shown in Fig. 7) estimated in the six
1348 simulations as listed in Table 1. The CARES G-1 flight tracks below 1 km (black line)
1349 during June 2010 are also shown with the Mv20Noah result; the terrain height is also
1350 shown as the black contour lines with the Mv21V3 result. b) Same as a) except around
1351 the CalNex observational sites Bakersfield and Pasadena (the red box shown in Fig. 7).
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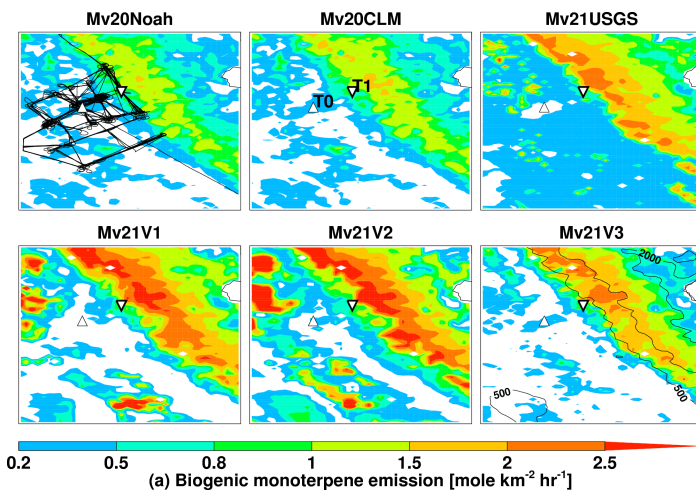
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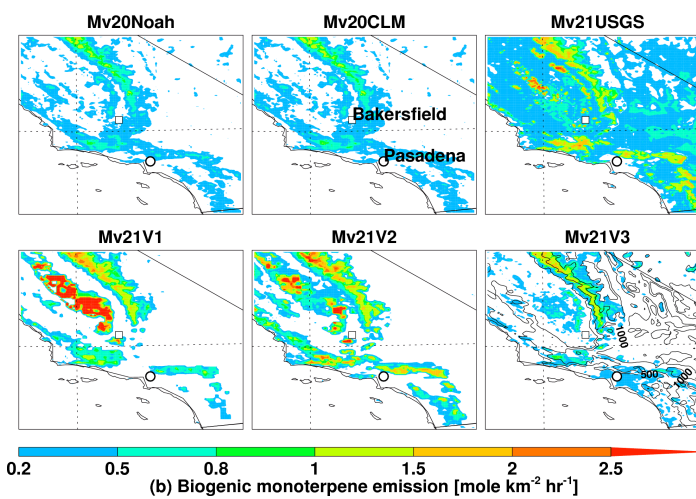
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1360 | **Figure 9.** Same as Fig. 8, except for biogenic monoterpene emissions.

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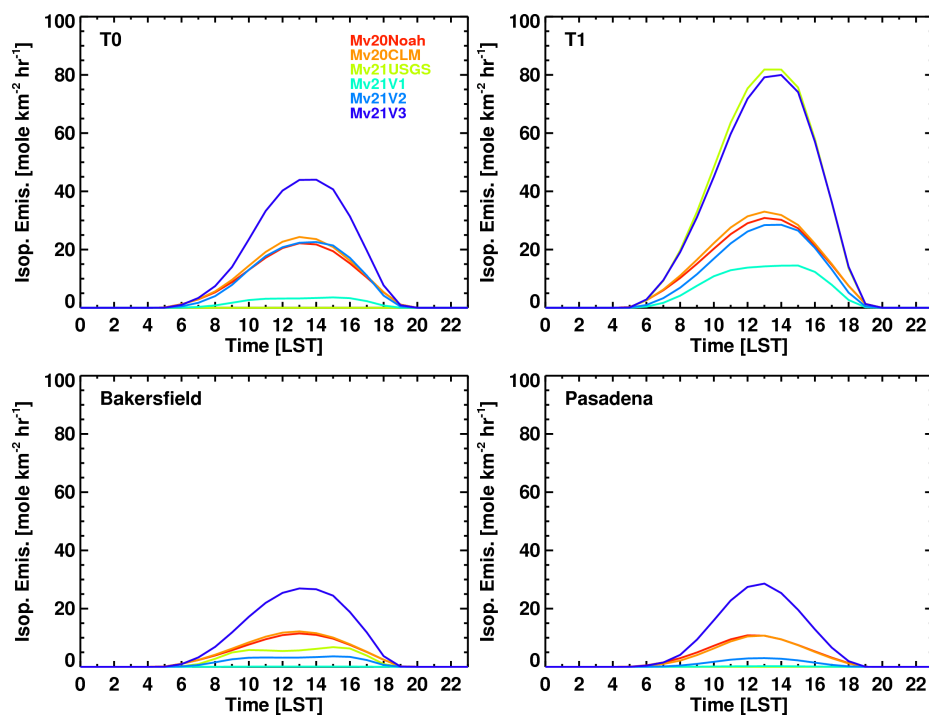
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1370 | **Figure 10.** Average diurnal variation of biogenic isoprene emissions at the four
1371 observation sites from the six simulations listed in Table 1.

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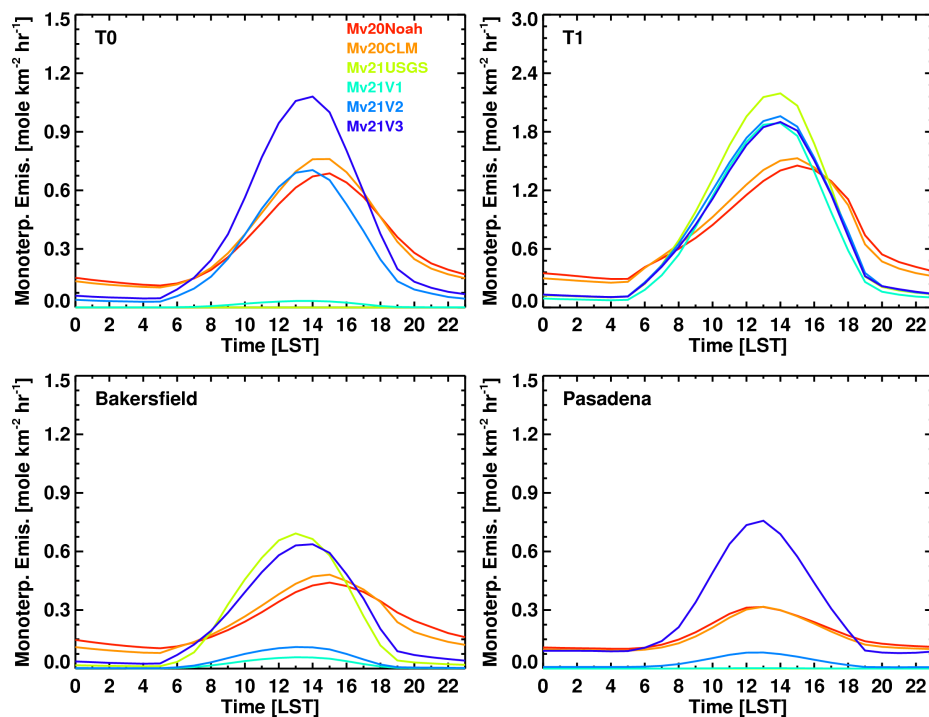
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Figure 11. Same as Fig. 10, except for biogenic monoterpene emissions.

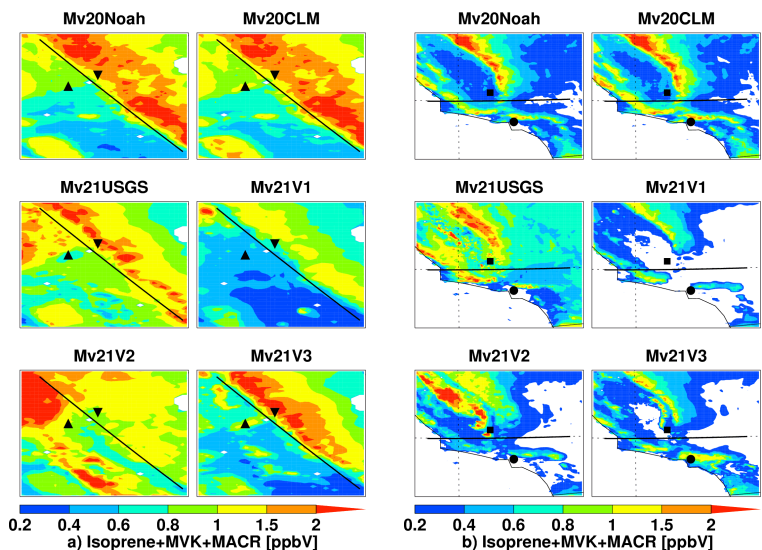
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1398 | **Figure 12.** a) Spatial distributions of monthly averaged surface isoprene mixing ratios
1399 | around the CARES T0 and T1 observational sites from the six simulations as listed in
1400 | Table 1. The black lines parallel to the Sierra Nevada divide the region to the Southwest
1401 | and the Northeast for comparison with CARES G-1 aircraft measurements shown in Fig.
1402 | 16 and 17. b) Same as a) except around the CalNex observational sites Bakersfield and
1403 | Pasadena. The black lines divide the region to southern California and the Central Valley
1404 | for comparison with CalNex WP-3D aircraft measurements shown in Fig. 16 and 17.

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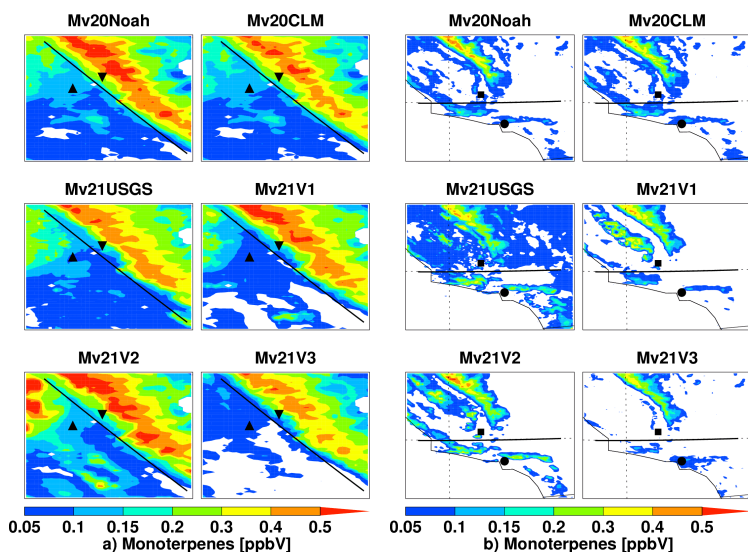
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1421 | **Figure 13.** Same as Fig. 12, except for monoterpene.

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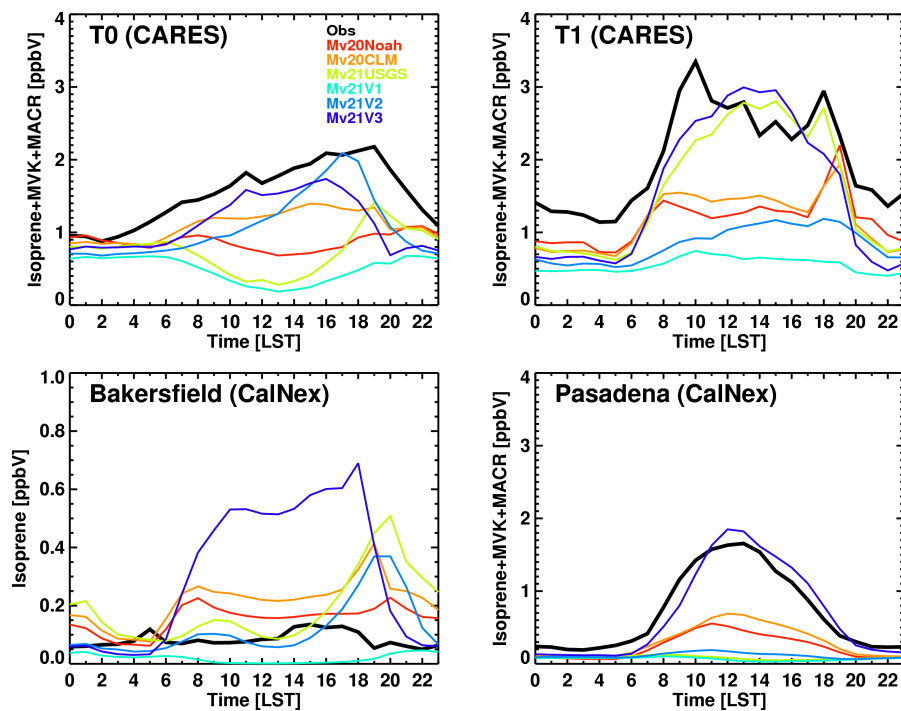
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Figure 14. Monthly averaged diurnal variation of surface isoprene+MVK+MACR mixing ratios at the three observation sites and isoprene mixing ratios at the Bakersfield site from the observations and six simulations listed in Table 2. The simulated values for the Bakersfield and Pasadena sites are averaged for the first two weeks of June to be consistent with the observations.

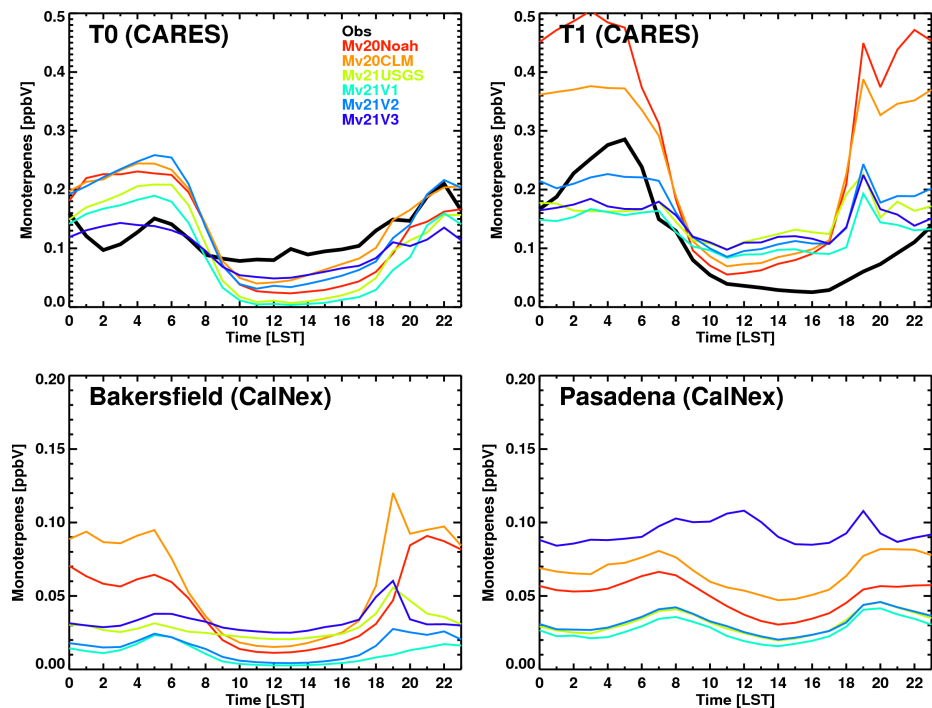
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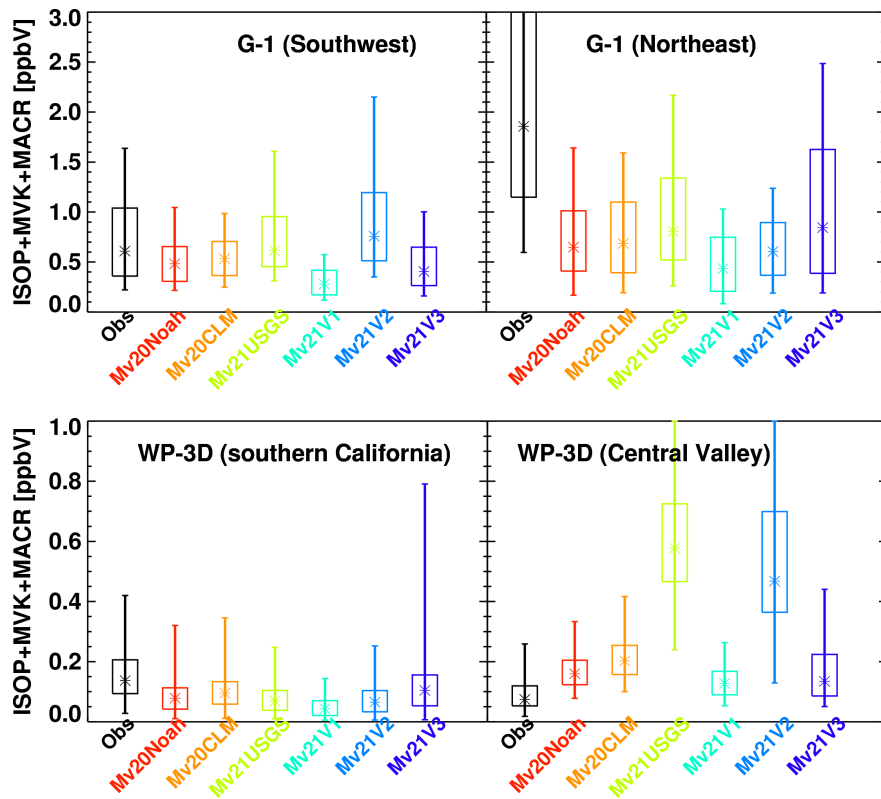


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Figure 15. Monthly averaged diurnal variation of surface monoterpene mixing ratios at the four observation sites from the observations and six simulations as listed in Table 2. There are no observations for the Bakersfield and Pasadena sites in June.

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1467 | **Figure 16.** Comparison of isoprene+MVK+MACR mixing ratios averaged below 1 km
1468 | from the observations by G-1 flights over the Southwest and Northeast regions (as
1469 | marked in Fig. 12a) and WP-3D flights over southern California and the Central Valley
1470 | (as marked in Fig. 12b) and the corresponding simulations. Asterisk denotes the 50th
1471 | percentiles. Vertical lines denote 10th and 90th percentiles, and the boxes denote the 25th
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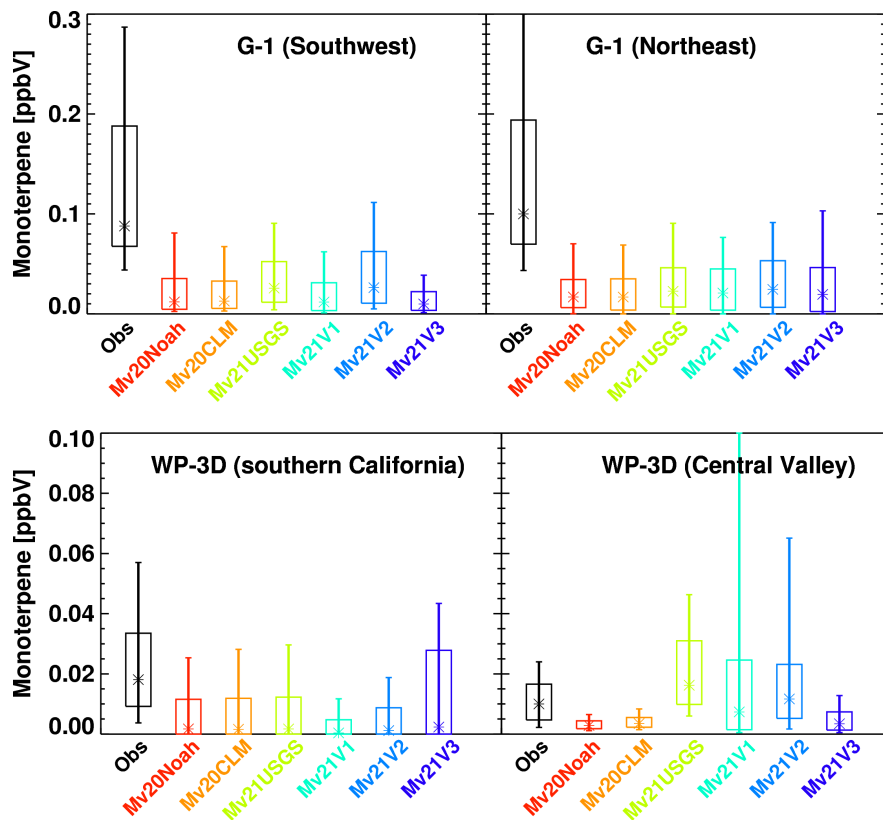
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Figure 17. Same as Fig. 16 except for monoterpane mixing ratios.

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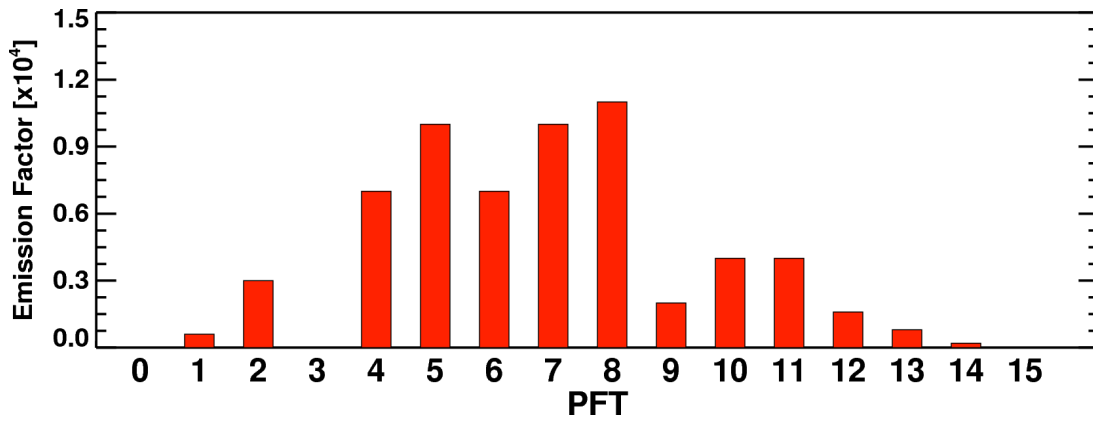


Figure 3. Biogenic isoprene emission factor for each PFT in MEGAN v2.1. The PFT number

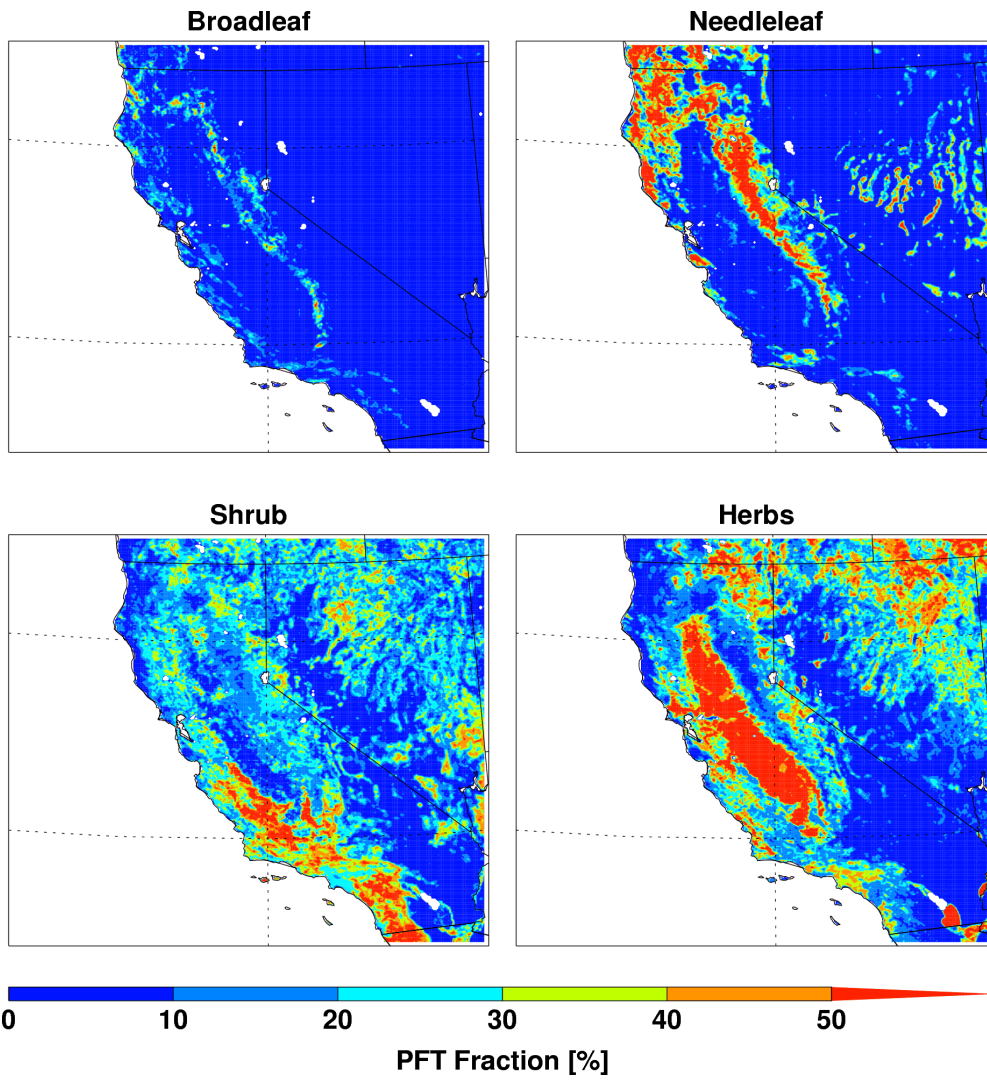


Figure 4.

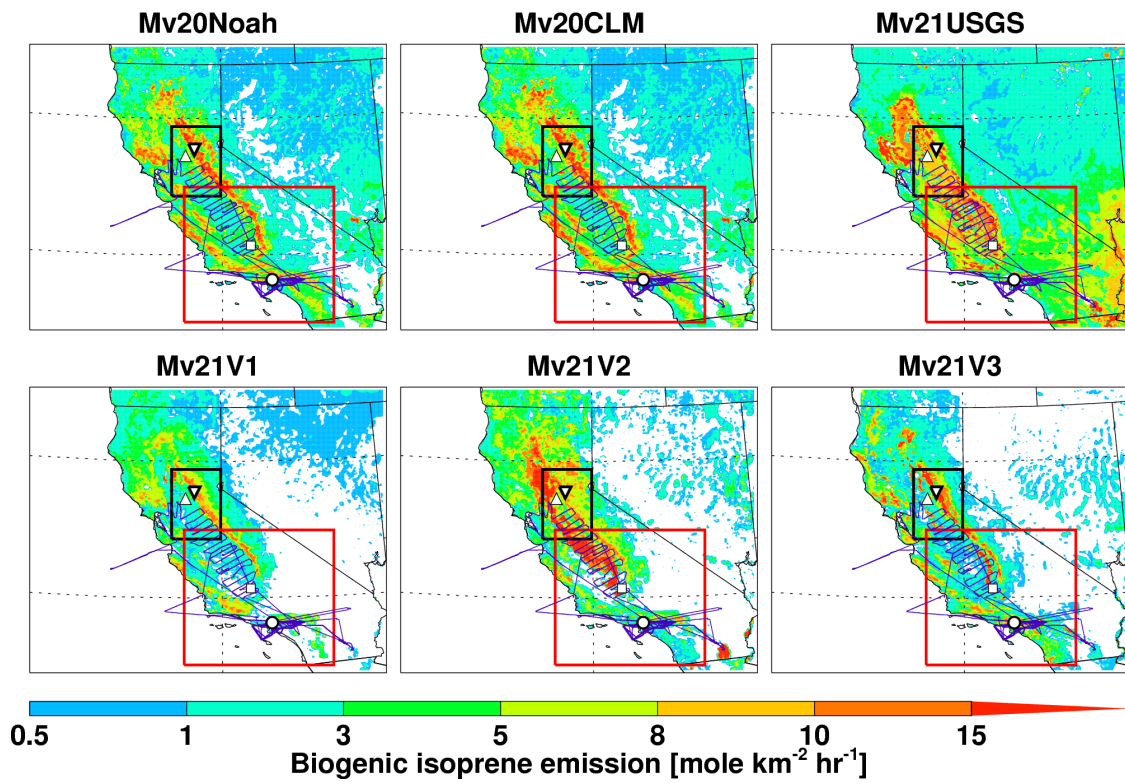


Figure 5.