The authors would like to thank both reviewers for their constructive comments. The manuscript has been updated to address the specific concerns raised. A brief response to the individual comments follows.

Reviewer 1

p. 2: Aromatic species were named to react with OH, NO3 and chlorine radicals only. This is true for all the explicitly named and most of the aromatic species except styrene, which may react with ozone too (Atkinson et al., 2006). Although being a minor contribution to the total petrol vapour concentration usually observed the SOA yield by styrene + ozone will be substantially larger as ozone reactions yield substantially less volatile species (Hallquist et al., 2009).

The chamber experiments on which the model is based were only performed with toluene and *m*-xylene, and thus inclusion of ozone reactions with styrene are beyond the scope of this work. However, two sentences were added to the final paragraph of section 2.4 pointing out this limitation with respect to styrene-derived SOA predictions.

p.5 (general): I would recommend providing a supporting only information document as several essential details of former studies are not directly available. This would help the reader understanding the study details much easier than to read in all the precessing articles. It includes a table of all the SOA groups and corresponding marker species (before and) after the modifications. An overview plot within the study may ease the understanding of the change when splitting B4. Parts of that Table can be found in Tables 1 and 2 as well as in Figure 1.

A table listing the SOA groups, their component species and surrogates before and after the modifications made as part of this work has been added as Supplemental Information.

A second aspect worth mentioning covers the initial conditions of the experiments displayed in Figs. 2, 3 and 5-8. Was an initial amount of seed aerosol introduced for partitioning calculations and if so which amount? Or was the system treated clean and the SOA had to form out of the gas-phase via new particle formation? This may cause a changing saturation vapour pressure because of Kelvin effects. The consequence would be time delays in experimental observations and challenges for simulation as partitioning requires an at least infinitesimal amount of pre-existing SOA mass. Was there a spin-up time assumed? Did the authors exclude certain data that were out of a certain range? In which concentration range of organic aerosol (OA) the authors would classify the approach to work properly and in which not?

It is important to note that no experiments were performed as part of this work. The updated model is based on the work of Xu et al. in which they simulated previously performed (and published) chamber experiments to evaluate a model for SOA formation from aromatic oxidation. One of the objectives of the current work is to

use the Xu et al. model to evaluate the impact of aromatic-derived SOA formation on regional air quality in an urban area. Thus, conditions typical to the South Coast Air Basin of California (SoCAB) were used as initial conditions in the model, which includes some background aerosol. New particle formation from gas-phase precursors typically occurs only when background aerosol concentrations are low, which is not usually the case in urban areas, or under very specific conditions not relevant to the simulated period.

p. 8: I would favour some "rapid writing style" improvements such as for "Ashworth et al. (2015) reported several updates to CACM (what they termed CACM0.0)... The rate constants for the reaction of organic peroxy radicals (RO2) with hydroperoxy radicals (HO2) or other RO2 species were increased to be in agreement with Table A2 in Ashworth et al. (2015).". Both rate constants were increased by which factor etc? Was the change substantial i.e. affecting any of the results displayed later on? Note Ashworth et al. (2015) did a forest study while this study deals with urban areas and does not include other biogenic VOCs than isoprene.

The organic peroxy radicals (RO2) are a class of species and include a number of individual species (RO21, RO22, RO23, etc.). Therefore, their cross reactions and reaction with HO<sub>2</sub> involve a large number of reactions, each with rate constants that have been updated according to Ashworth et al. (2015). The modifications related to biogenic oxidation chemistry are described in detail in Ashworth et al. (2015). The implementation here was performed with only the few modifications described in the manuscript. Readers interested in the specific details of their model updates are now referred to their original published work, including the supplemental information table SA2.

Inclusion of the modifications described in Ashworth (2015) had a negligible effect on the model results. This was expected because, as the reviewer states, their work was related to modeling the air above a forest canopy, and biogenics play a much less prominent role in the Los Angeles urban atmosphere. The paragraph describing the effects of their suggested modifications has been updated to make this clear.

p.9 (general): The redistribution of class B4 to subspecies and the related results left me somewhat puzzled when looking at Figure 3. While the saturation vapour pressure estimation had a notable (+1  $\mu$ g/m3) effect the daily pattern remained unchanged. Both rush-hours are visible at different intensity, which agrees with the different mixing layer height and dilution. This daily pattern is not obtained for the simulation of the redistributed B4 SOA class that basically shows an inverse daily structure of a mixing layer height and substantially lower formation yields. In typically elevated urban conditions I would expect to obtain similar results for a lumped compound class and for the individual compounds if using the Pankow (199a, 1994b) approach, since there is no condensation but partitioning taking place at any concentration level. My two only explanations for that would be (i) one of both approaches used a completely different saturation vapour pressure or (ii) the preexisting organic aerosol mass is

# subestimated with a negative feedback on the formation rate. Thus, I don't know which of both simulations to trust as no observations are displayed for quality check.

The authors agree with the reviewer's explanations. Splitting SOA group B4 into three 'groups' with only one species each had the effect of changing the saturation vapor pressure for the two species that were not the surrogate in the original group B4. In the updated model, their partitioning is calculated based on their own structures using the SIMPOL.1 method, instead of on the structure of the original surrogate (AP12). Thus the updated model can be expected to be more accurate, although there are admittedly no observations with which to compare the model predictions. In general, the Pankow approach using lumped species does a good job at predicting SOA partitioning. However, what we believe our results show is that when species with similar structures, but high concentrations relative to other SOA species, are lumped into the same SOA group, this can have dramatic effects on modeled SOA.

One explanation for the change in diurnal profile of SOA concentration is that the long-chain functionalized alkanes included in the original SOA group B4 (AP11, AP12 and UR20) are strongly correlated with vehicle emissions. Splitting SOA group B4 into groups B4, B6 and B7 results in less contribution to SOA from AP11 (B6) and UR20 (B7) due to their higher SIMPOL.1-modeled saturation vapor pressure. Thus, other SOA species, which may form slowly in the atmosphere, or have sources other than vehicle emissions, contribute more strongly to the net SOA concentration profile. It should also be noted that the time profiles shown in Fig.3 are domain averaged. It is likely that this profile may look different for specific locations within the domain (e.g., downtown Los Angeles).

For a better understanding I would recommend a replacement of the class numbers like A1 by a structure based name in Fig. 3, as the abbreviations have not been explained all in the text and one could easier identify the responsible group for the deviations.

Unlike the new aromatic-derived lumped SOA species, the lumped SOA species developed prior to this work do not lend themselves to a simple structure-based classification, particularly after the re-lumping reported here. A detailed structure-based name for each group would 1) primarily represent only the surrogate species, and 2) be too long for easy inclusion in a figure. However, a reference has been added to the caption for Figure 3 referring the reader to Figure 4 for the surrogate structures for each of the lumped SOA species.

p.11: It makes me struggle somewhat that the authors used nicely the SIMPOL.1 approach (Pankow and Asher, 2008) for a better saturation vapour pressure estimation but adjust it afterwards to match the former results by Xu et al. (2015) at a temperature of 25 degrees C. Did the authors not trust the SIMPOL.1 estimates? Was the approach used to get the relative list (volatile, semi-volatile, non-volatile compounds) in the correct shape but control the absolute values? Please explain. This adjustment was performed just as an "extreme-case" scenario to evaluate the relative importance of aqueous- and organic-phase partitioning of the aromatic-derived SOA species. It is not included in the final version of the model. The paragraph describing this test has been updated to make this more clear.

General: A key feature for every experiment and simulation approach - not this one exclusively - should be a statement in which range a certain approach provides reasonable estimates (valid range). This would force future users to carefully consider not a "black box" for application and take it into account for interpretation of results. Could the authors provide such to make a future application as appropriate as possible and potentially name issues with a need for further improvement? That could serve as new standard.

The range of experimental concentrations used for model development are described in Xu et al. (2015).

Finally one question about other SOA precursor and compounds was left: So far larger biogenic VOCs such as monoterpenes are not included probably because of simulation capacities and lack of information. Those are much less volatile and would allow a higher preexisting mass present for partitioning of aromatic compounds. Did the authors made tests on the sensitivity of the simulations to preexisting particle mass?

It should be noted that monoterpenes and isoprene are included in the reaction mechanism. The lack of model sensitivity to the modifications of Ashworth et al (2015) is evidence that biogenics play a smaller role in the SoCAB region compared to a forest canopy; in addition, the modifications are likely to be relevant in low-NO<sub>X</sub> scenarios not representative of the SoCAB. Finally, background aerosol concentrations are generally high in this region. Thus, the sensitivities described by the reviewer were not explored, as the focus of this work was specifically SOA formation from aromatic oxidation.

p. 24 and 26 (Fig. 6 and 8): I guess the apparent notable presence of furanones and tiny amounts of epoxides in Fig. 6 can be explained by the primary focus on AROH emissions. Correct?

This follows closely the composition of SOA formed in the chamber simulations of Xu et al (2015). In the CIT model, the AROH emissions were used to estimate the emissions of toluene and m-xylene in the SoCAB region. Thirty percent of the original AROH emissions were changed to toluene, another 30% to m-xylene, and the remaining 40% was left as AROH. The AROH species followed the original oxidation pathways in CACM, and would therefore not contribute to the new SOA groups (C1-C5).

p. 26 (Fig. 8): The daily structure of PAN-like species and epoxides is very interesting and matches with expectations. Could you provide a standard deviation (variation) of your mean model domain pattern for the 5 classes?

A table with average concentration and one standard deviation in concentration for the five aromatic-derived SOA species has been added to the Supplementary information. Manuscript prepared for Geosci. Model Dev. with version 2015/04/24 7.83 Copernicus papers of the LATEX class copernicus.cls. Date: 7 April 2016

# **Development of aroCACM/MPMPO 1.0: A Model to Simulate Secondary Organic Aerosol from Aromatic Precursors in Regional Models**

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Abstract. The atmospheric oxidation of aromatic compounds is an important source of secondary organic aerosol (SOA) in urban areas. The oxidation of aromatics depends strongly on the levels of nitrogen oxides ( $NO_X$ ). However, details of the mechanisms by which oxidation occurs are only recently being elucidated. Xu et al. (2015) developed an updated version of the gas-phase Caltech

- 5 Atmospheric Chemistry Mechanism (CACM) designed to simulate toluene and *m*-xylene oxidation in chamber experiments over a range of  $NO_X$  conditions. The output from such a mechanism can be used in thermodynamic predictions of gas-particle partitioning leading to SOA. The current work reports the development of a model for SOA formation that combines the gas-phase mechanism of Xu et al. (2015) with an updated lumped SOA partitioning scheme (MPMPO) that allows partitioning
- 10 to multiple aerosol phases and that is designed for use in larger scale three-dimensional models. The resulting model is termed aroCACM/MPMPO 1.0. The model is integrated into the University of California, Irvine California Institute of Technology (UCI-CIT) airshed model, which simulates the South Coast Air Basin (SoCAB) of California. Simulations using 2012 emissions indicate that "low-NO<sub>X</sub>" pathways to SOA formation from aromatic oxidation play an important role, even in
- 15 regions that typically exhibit high  $NO_X$  concentrations.

#### 1 Introduction

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Atmospheric aerosol particles negatively affect human health, contribute to reduced visibility, and impact Earth's climate through their ability to scatter and absorb radiation and affect cloud properties (Finlayson-Pitts and Pitts, 2000; Seinfeld and Pandis, 2006). Secondary organic aerosol (SOA) accounts for a significant fraction of atmospheric aerosol mass (Hallquist et al., 2009; Kanakidou et al., 2005; Krall and Sainfeld, 2008). In other areas, the gas phase aridation of atmospheric is

2005; Kroll and Seinfeld, 2008). In urban areas, the gas-phase oxidation of aromatic hydrocarbons is thought to be an important contributor to SOA formation, although the fundamental mechanisms are only recently being elucidated (Im et al., 2014; Ng et al., 2007; Song et al., 2005; Xu et al., 2015).

In urban areas, aromatics are an important constituent of the mix of volatile organic compounds

- 25 (VOCs) and are emitted primarily from industrial operations, fuel evaporation, and vehicle exhaust, with smaller contributions from biomass burning and other sources (Karl et al., 2009). In addition, there is evidence of biogenic sources of aromatic compounds that may be important in rural areas (Gratien et al., 2011; White et al., 2009).
- The atmospheric oxidation of aromatics is thought to proceed primarily by reaction with hydroxyl
  radical (OH), but nitrate (NO<sub>3</sub>) and chlorine (Cl) radicals readily oxidize some substituted aromatics as well (Finlayson-Pitts and Pitts, 2000). The oxidation pathway (addition to the ring as opposed to hydrogen atom extraction from a side-chain) depends on the oxidant and the extent of substitution. Once initiated, oxidation proceeds through a variety of pathways that depend strongly on the level of nitrogen oxides (NO<sub>X</sub>) that are present. In most cases, SOA yields from aromatics are negatively
  correlated with NO<sub>X</sub> concentrations (Ng et al., 2007; Song et al., 2005; Xu et al., 2015), though
- exceptions have been reported (Im et al., 2014).

In order to simulate the formation of SOA and other secondary pollutants such as ozone, air quality models require chemical mechanisms that predict gas-phase oxidation chemistry over a wide range of  $NO_X$  concentrations in a computationally efficient manner. Such mechanisms can be highly

- 40 reduced ("lumped") or highly specific depending on computational demands and application. For example, the computationally efficient mechanism of Carter et al. (2012) uses non-specific reactions to simulate the formation of secondary products, but does not track oxidation products specifically. On the other hand, Im et al. (2014) developed the comprehensive UNIPAR model to predict SOA formation from toluene and 1,2,3-trimethylbenzene. UNIPAR uses the detailed Master Chemical
- 45 Mechanism (MCM (Jenkin et al., 2003; Saunders et al., 2003)) to simulate aromatic SOA formation based on a lumped equilibrium-partitioning scheme that accounts for liquid-liquid phase separation, aqueous-phase oligomerization, and organo-sulfate formation. The UNIPAR output in conjunction with outdoor chamber experiments performed over a range of NO<sub>X</sub> conditions suggest an important role for aerosol-phase oligomerization in aromatic SOA formation.
- 50 As an intermediate approach to the non-specific and comprehensive methods mentioned above, Xu et al. (2015) updated the Caltech Atmospheric Chemical Mechanism (CACM) to include SOA formation from the gas-phase oxidation of toluene and *m*-xylene and evaluated the model through comparison to chamber experiments. When combined with a partitioning model based on absorption theory (Pankow, 1994), the updated CACM simulated SOA formation accurately over a range of
- 55 NO<sub>X</sub> conditions, without the need for separate low- and high- NO<sub>X</sub> parameters. In addition, the reduced complexity of CACM compared to the fully explicit MCM decreases computational cost, making it ideal for incorporation into larger-scale models.

This work reports results from the development of an updated SOA-formation mechanism, aro-CACM/MPMPO 1.0, based on the initial work of Xu et al. (2015) and an updated version of the Model to Predict the Multi-phase Partitioning of Organics (MPMPO, (Griffin et al., 2003)). The

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testbed for this mechanism is the University of California, Irvine – California Institute of Technology (UCI-CIT) Airshed model, which simulates the South Coast Air Basin (SoCAB) of California. This work is presented in two parts, each with the aim of improving SOA predictions in regional models covering a range of NO<sub>X</sub> conditions. First, the vapor pressure calculations for all lumped

65 SOA species in the model were updated using the SIMPOL.1 method of Pankow and Asher (2008), and lumped SOA species were redistributed based on thermodynamic properties and relative abundance. Second, the updated CACM of Xu et al. (2015) was incorporated along with appropriate lumped SOA species and a treatment of aqueous-phase partitioning of aromatic-derived SOA.

## 2 Model Description

### 70 2.1 UCI-CIT Airshed Model

The UCI-CIT model is an Eulerian chemical transport model that solves simultaneously the advection/diffusion equation coupled with emission, deposition, and chemistry terms. The model domain is divided horizontally into an 80 x 30 grid of 5 km x 5 km cells and vertically into five layers reaching a height of 1100 m above the surface. The domain includes the entire SoCAB, which includes

- 75 Orange County and parts of Los Angeles, Ventura, San Bernardino and Riverside Counties. The advection/diffusion equation is solved using the Quintic Splines Taylor Series Expansion approach described by Nguyen and Dabdub (2001). Deposition is based on the dry deposition flux described by Wesely for gas-phase species (Wesely, 1989), and a combination of diffusional deposition and gravitational settling for particles (Griffin et al., 2002a).
- 80 The current version of the UCI-CIT model utilizes the previous versions of CACM and MPMPO to describe formation of secondary pollutants, including SOA (Griffin et al., 2005, 2002b, a; Pun et al., 2002). Advanced numerical algorithms are used to solve the non-linear system of highly-coupled differential equations involved in a parallel computational architecture (Dabdub and Seinfeld, 1996, 1994; Nguyen and Dabdub, 2002, 2001).

#### 85 2.2 Emissions and Meteorology

Evaluation of aroCACM/MPMPO 1.0 described subsequently is performed with a three-day simulation using meteorological data representative of the meteorology in the SoCAB (Carreras-Sospedra et al., 2010). It is important to note that in the testing and development presented here, a specific historical event is not being simulated.

90 As part of this work, the UCI-CIT model was updated to accept emissions based on the 2012 Air Quality Management Plan (AQMP) provided by the South Coast Air Quality Management District (SCAQMD, 2013), which include current (2008) emissions. All plots shown are for the second simulation day with 2008 emissions.

#### 2.3 Updated Vapor Pressure Calculations and New SOA Lumping Scheme

- 95 In the UCI-CIT model, the most recent version of MPMPO is used to calculate gas-particle conversion of secondary organic species (Griffin et al., 2005, 2003). In species-specific equilibriumpartitioning based models, including MPMPO, thermodynamic properties including sub-cooled liquid vapor pressure ( $p_L^O$ ) are required to determine the extent of gas-particle conversion (Pankow, 1994). However, experimental values for  $p_L^O$  are often unavailable, and methods for estimating  $p_L^O$
- 100 based on molecular structure and/or other properties are required. The SIMPOL.1 group-contribution method of Pankow and Asher (2008) is adopted here to calculate updated SOA species vapor pressures for use in MPMPO. This method was selected because it includes contributions from functional groups that are present in the SOA species used in MPMPO but were not able to be included using the previously utilized method to calculate  $p_L^O$  (Myrdal and Yalkowsky, 1997). In addition, SIM-
- 105 POL.1 was parameterized and tested using a large set of experimental  $p_L^O$  data that spans 14 orders of magnitude (Pankow and Asher, 2008).

Due to computational limitations, individual SOA species of similar molecular structure are often 'lumped' into groups for the purpose of calculating gas-particle conversion in equilibriumpartitioning based models when applied in three-dimensional air quality models. Pun et al. (2002)

- 110 developed the original SOA lumping scheme used in the UCI-CIT model. Further modifications to the scheme were made by Griffin et al. (2005). Following the inclusion of SIMPOL.1 for  $p_L^O$  calculations, the existing SOA lumping scheme used by MPMPO was re-evaluated based on the structure, calculated thermodynamic parameters, and relative abundance of the individual (i.e. non-lumped) SOA species. The motivation behind the specific changes made as part of this work, and their effect 115 on model output are described in Section 3.1.
  - 2.4 Aromatic-Derived SOA Formation

The gas-phase mechanism for toluene and *m*-xylene oxidation under varying  $NO_X$  conditions originally developed by Xu et al. (2015) was incorporated into the UCI-CIT model without modification. However, gas-particle conversion for the new aromatic-derived SOA species has been adapted for

- 120 inclusion into a large-scale three-dimensional model. Because of computational limitations in such models, an SOA lumping scheme was applied to the aromatic-derived SOA species, as in the previous version of MPMPO. The new species are divided into five SOA groups labeled C1 C5. Table 1 describes the new SOA groups and lists the species and surrogates for each group. Structures for these surrogate species are shown in Figure 1. Individual species included in the new SOA groups
- 125 were selected based on structure and relative yields in the zero-dimensional model simulations of Xu et al. (2015).

As with the pre-existing SOA species,  $p_L^O$  for the new surrogate SOA species are calculated using the SIMPOL.1 method described in section 2.3. These new, highly oxygenated SOA groups also could be expected to partition favorably into the aqueous phase. The MPMPO is a fully coupled

- 130 mixed-phase SOA model that accounts for partitioning of each SOA group into an aqueous and organic aerosol phase simultaneously (Griffin et al., 2005, 2003). Therefore, Henry's Law constants (HLC) for the new species were calculated using the method of Suzuki et al. (1992) for use in the aqueous-phase partitioning module of MPMPO. Finally, due to the new grouping scheme, the recalculated  $p_L^O$ , and the introduction of aqueous-phase partitioning for the new SOA species, the
- 135 experimentally determined correction factors for  $p_L^O$  used by Xu et al. (2015) were not adopted in aroCACM/MPMPO. This could result in an underestimation of SOA formation. To explore further this potential underestimation of SOA, several simulations were performed in which correction factors were applied to bring the SIMPOL.1-calculated  $p_L^O$  into agreement with those used by Xu et al. (2015) at 298 K. Results from these simulations are discussed in section 3.2.
- 140 Emissions of aromatic compounds in the 2012 AQMP are grouped into high-SOA-yield (AROH) and low-SOA-yield (AROL) species. In the absence of more detailed speciation data, some estimation of the contribution of toluene and *m*-xylene to total aromatics must be made. To evaluate the impact of toluene and *m*-xylene oxidation on SOA concentrations individually, model runs are presented in which AROH emissions are assumed to be entirely toluene or entirely *m*-xylene. A
- 145 scenario where toluene and *m*-xylene are assumed to be 30% each of AROH is also presented, as an approximation of the impact of this new chemistry in a real urban environment. Although this is a rough estimate, toluene and *m*-xylene are known to be major contributors to aromatic emissions in urban areas (de Gouw et al., 2005), and other aromatics, including *o* and *p*-xylene, likely exhibit similar reactivity. Some aromatic compounds that are present in the atmosphere, such as styrene,
- 150 are expected to have higher SOA yields due to their reaction with ozone (Hallquist et al., 2009). However, because of their relatively low concentrations, these species are not specifically treated in this work. Unless otherwise noted, all results shown for model simulations that include the new aromatic-derived SOA species use this final emissions scenario that includes *m*-xylene and toluene.

#### 2.5 Low-NO<sub>X</sub> Chemistry

- 155 Ashworth et al. (2015) reported several updates to CACM (what they termed CACM0.0), primarily related to low-NO<sub>X</sub> chemistry, based on a comparison to field measurements from a forest canopy and to other gas-phase chemical mechanisms. While specific monoterpene oxidation chemistry for  $\alpha$ -pinene,  $\beta$ -pinene, and limonene is not included in the UCI-CIT version of CACM, two of their recommended adjustments were made. The rate constants for the reaction of organic peroxy radicals
- 160 (RO2) with hydroperoxy radicals (HO2) or other RO2 species were increased to be in agreement with Table A2 in Ashworth et al. (2015). Additionally, the formation of isoprene-derived organic nitrates, along with their subsequent reactions with OH, were added. However, these modifications had little effect on modeled NO and NO<sub>2</sub> concentrations or SOA concentration<del>, as This</del> would be

expected in this as biogenics play a much greater role in SOA formation in a forested region (for

165 which their model was developed), than in the relatively high-NO<sub>X</sub>arid area, arid SoCAB region.

#### **3** Results and Discussion

#### 3.1 Updated Vapor Pressure and New SOA Lumping Scheme

The change in predicted SOA upon implementation of the SIMPOL.1 method for vapor pressure calculations is shown in Figures 2 and 3. Use of the SIMPOL.1 vapor pressure calculations results in
a modest increase in SOA concentration over the base case (Fig. 2a,b). As seen in Figure 3(a,b), this increase can be attributed primarily to SOA group B4, which includes oxidation products of long-chain alkanes (Griffin et al., 2005; Pun et al., 2002). Because of the large contribution of SOA group B4 to total SOA concentration, a redistribution of SOA species was considered and performed. A summary of modifications to the SOA lumping scheme is presented in Table 2. The structures of the

175 resulting 12 surrogate SOA species are shown in Figure 4.

The new redistribution results in a large decrease in total SOA concentration (Fig. 2c). The contribution of each SOA group to total SOA in the redistributed scheme is shown in Figure 3c. Splitting SOA group B4 (long-chain alkane oxidation products) into groups B4, B6 and B7 (Table 2) results in an order-of-magnitude decrease in SOA from these species. Previously, the surrogate species for

- 180 group B4 had both the lowest modeled vapor pressure and lowest domain-wide average concentration of all species in that SOA group. This had resulted in an over-estimation of SOA from this source, as can be seen by comparing Figures 3b and 3c. The other major effect of the redistribution of SOA is the decrease in SOA formation from group A2. Although the use of species UR3 as the surrogate for SOA group A2 results in a lower modeled  $p_L^O$  for this group, the largest contribution to
- 185 group A2 had come from species UR26, which has a higher  $p_L^O$  than the other species in group A2. Moving UR26 to group A4 and using it as the surrogate for this group results in a decrease in the combined SOA concentrations from groups A2 and A4.

#### 3.2 SOA Formation from Toluene and *m*-xylene Oxidation

- Figure 5 shows 24-hr average SOA concentrations assuming AROH emissions are all *m*-xylene, all
  toluene, or 30% each *m*-xylene and toluene (with the remaining 40% following the original CACM AROH chemistry). Comparison with Figure 6c shows that aromatic-derived SOA accounts for 75% or more of total modeled SOA concentration in the domain. In these simulations, *m*-xylene oxidation leads to higher total SOA concentration than does toluene oxidation (Fig. 5a,b). The opposite trend is seen in the simulation results of Xu et al. (2015) and the experiments on which the model
- 195 was based, both of which were performed under dry conditions at room temperature. Several factors may contribute to this difference in results. First, vapor pressure correction factors used by Xu et al. (2015) were not used here. In addition, the partitioning model here allows for aqueous-phase SOA

formation. Finally, it should be noted that the previous experiments and simulations allowed the parent hydrocarbon to react almost to completion. In fact, the reaction of m-xylene with OH occurs

200 more rapidly than does that of toluene, such that a larger fraction of *m*-xylene reacts in that simulation, potentially leading to more SOA formation despite having a smaller yield when observed in the laboratory.

As shown in Figure 6 for the *m*-xylene and toluene scenarios, the speciated SOA compositions are in reasonable agreement with the results of Xu et al. (2015) under high-NO<sub>X</sub> conditions. In both

cases, there are strong contributions from furanones (C1) and peroxy-nitrates (C3), and in the case of toluene, peroxy-bicyclic rings (C5). However, in both the *m*-xylene and toluene cases, the relative SOA contribution from phenols (C2) is smaller than would be expected from the zero-dimensional model simulations of Xu et al. (2015) for high-NO<sub>X</sub> conditions and corresponds more closely to low-NO<sub>X</sub> conditions. It is important to note that these conditions are total NO<sub>X</sub> and do not generally
consider the split of NO<sub>X</sub> between NO and NO<sub>2</sub>.

Because the  $NO_X$  levels in the SoCAB are relatively high regardless of time of day, modeled domain average NO and  $NO_2$  concentrations are shown in Figure 7a. In general, the contribution to total SOA from groups C1, C2 and C4 are correlated with high NO concentrations, and groups C3 and C5 with lower relative NO concentrations (Fig. 7b) and peak OH concentrations (early

- 215 afternoon; not shown). In the gas-phase aromatic oxidation mechanism, NO competes with NO<sub>2</sub> to react with acyl peroxy radical species. When the [NO]/[NO<sub>2</sub>] ratio is low, the NO<sub>2</sub> reaction forms peroxy-nitrates and leads to an increase in group C3. When the [NO]/[NO<sub>2</sub>] ratio is high, the NO reaction with certain peroxy radical species can form furanones and lead to an increase in group C1. The reaction of NO<sub>2</sub> with phenolic radical species to form nitro-phenols leads to an increase
- 220 in SOA formation from group C2 when NO<sub>2</sub> concentrations are highest. Meanwhile, NO competes with the intra-molecular cyclization reaction of certain peroxy radical species, leading to a decrease in peroxy bicyclic ring species (SOA group C5) under high-NO conditions. The diurnal trend of SOA derived from group C4 is more difficult to explain based on the gas-phase mechanism. Some contribution to the decrease seen in groups C1, C2 and C4 may be due to the increased preference
- for the gas phase of these semi-volatile species during the daytime, as is supported by the average fraction of species C1, C2 and C4 in the particle phase as a function of time of day (Fig. 8).

Although the experimental validation of the mechanism used in the current work was for dry conditions, aqueous-phase partitioning of aromatic-derived SOA species was incorporated into aro-CACM/MPMPO 1.0. To evaluate the importance of aqueous-phase partitioning for the aromatic-

230 derived SOA species in the updated UCI-CIT model, a simulation was performed in which species C1 – C5 were allowed to partition only to the organic phase. Results indicate that > 99% of modeled SOA from species C1 – C5 is in the aqueous-phase. It is possible that the vapor pressure correction factors used by Xu et al. (2015) would increase the fraction of species C1 – C5 in the organic phase. Therefore, a final simulation was performed in which the SIMPOL.1-calculated vapor pressure

- sures were adjusted to match those used by Xu et al. (2015) at 298K. (Note: this model run includes aqueous-phase partitioning of species C1 C5.) This results in little change to modeled SOA, as seen by comparing Figures 5c and 5d, again suggesting that aqueous-phase partitioning may play a dominant role in SOA formation from the aromatic-derived SOA species. (Note: In the final version of the model, the unadjusted SIMPOL.1 vapor pressures are used.)
- In addition to aqueous-phase partitioning, results from other experimental and modeling work suggest that aqueous-phase oligomerization may be an important route to SOA formation from aromatics under atmospherically relevant conditions (Im et al., 2014). Future studies using experiments on aromatic oxidation under a variety of relative humidity conditions will be required to assess further the relative importance of aqueous- vs. organic-phase partitioning and the role of oligimerization and to refine the aroCACM/MPMPO 1.0 model.

#### 4 Conclusions

This work reports the development of aroCACM/MPMPO, a gas-phase and SOA model designed for use in large-scale chemical transport models that includes important routes to SOA formation from the oxidation of aromatic species. aroCACM/MPMPO makes use of updated schemes for va-

- 250 por pressure estimation and grouping SOA species and includes both aqueous- and organic-phase partitioning for all SOA species using an equilibrium-partitioning module. Results presented here from the implementation of aroCACM/MPMPO 1.0 into the UCI-CIT Airshed model support the importance of aromatic-derived SOA in urban areas. Of particular note is the manner in which the concentrations of individual SOA-partitioning species demonstrate different trends with the time of
- 255 day. Furanones, phenols, and epoxides are found to increase their contribution to total SOA during times of high  $NO_X$  and low OH, while peroxy-nitrate and peroxy-bicyclic ring species show the opposite trend. This indicates that even in areas generally considered "high"  $NO_X$  it is important to include the contribution of low  $NO_X$  pathways to SOA formation from aromatic oxidation, primarily due to the importance of the NO to  $NO_X$  ratio.

#### 260 5 Code Availability

Model code is available by request.

Acknowledgements. This work was performed under grants from the National Science Foundation (CHE-1213623 and ATM-0901580). Any opinions, findings, and conclusions or recommendations expressed in this material are those of the authors and do not necessarily reflect the views of the National Science Foundation.

#### 265 References

275

- Ashworth, K., Chung, S. H., Griffin, R. J., Chen, J., Forkel, R., Bryan, A. M., and Steiner, A. L.: FORest canopy atmosphere transfer (FORCAsT) 1.0: a 1-D model of biosphere–atmosphere chemical exchange, Geosci. Model Dev. Discuss., 8, 5183–5234, 2015.
- Carreras-Sospedra, M., Vutukuru, S., Brouwer, J., and Dabdub, D.: Central power generation versus distributed
- 270 generation An air quality assessment in the South Coast Air Basin of California, Atmos. Environ., 44, 3215–3223, 2010.
  - Carter, W., Heo, G., Cocker, D., and Nakao, S.: SOA formation: Chamber study and model development. Final report to the California Air Resources Board, Contract No. 08-326, Tech. rep., 2012.
  - Dabdub, D. and Seinfeld, J. H.: Air quality modeling on massively parallel computers, Atmos. Environ., 28, 1679–1687, 1994.
  - Dabdub, D. and Seinfeld, J. H.: Parallel computation in atmospheric chemical modeling, Parallel Comput., 22, 111–130, 1996.
    - de Gouw, J. A., Middlebrook, A. M., Warneke, C., Goldan, P. D., Kuster, W. C., Roberts, J. M., Fehsenfeld, F. C., Worsnop, D. R., Canagaratna, M. R., Pszenny, A. a. P., Keene, W. C., Marchewka, M., Bertman, S. B.,
- 280 and Bates, T. S.: Budget of organic carbon in a polluted atmosphere: Results from the New England Air Quality Study in 2002, J. Geophys. Res., 110, D16 305, doi:10.1029/2004JD005623, 2005.

Finlayson-Pitts, B. and Pitts, J.: Chemistry of the upper and lower atmosphere – Theory, experiments and applications, Academic Press, San Diego, 2000.

- Gratien, A., Johnson, S. N., Ezell, M. J., Dawson, M. L., Bennett, R., and Finlayson-Pitts, B. J.: Surprising
- formation of *p*-cymene in the oxidation of  $\alpha$ -pinene in air by the atmospheric oxidants OH, O<sub>3</sub>, and NO<sub>3</sub>, Environ. Sci. Technol., 45, 2755–2760, 2011.
  - Griffin, R. J., Dabdub, D., Kleeman, M. J., Fraser, M. P., Cass, G. R., and Seinfeld, J. H.: Secondary organic aerosol 3. Urban/regional scale model of size- and composition-resolved aerosols, J. Geophys. Res. Atmos., 107, 4334, doi:10.1029/2001JD000544, 2002a.
- 290 Griffin, R. J., Dabdub, D., and Seinfeld, J. H.: Secondary organic aerosol 1. Atmospheric chemical mechanism for production of molecular constituents, J. Geophys. Res., 107, 4332, doi:10.1029/2001JD000541, 2002b.
  - Griffin, R. J., Nguyen, K., Dabdub, D., and Seinfeld, J. H.: A coupled hydrophobic-hydrophilic model for predicting secondary organic aerosol formation, J. Atmos. Chem., 44, 171–190, 2003.
  - Griffin, R. J., Dabdub, D., and Seinfeld, J. H.: Development and initial evaluation of a dynamic species-resolved
- 295 model for gas phase chemistry and size-resolved gas/particle partitioning associated with secondary organic aerosol formation, J. Geophys. Res., 110, D05 304, doi:10.1029/2004JD005219, 2005.
  - Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod,
- 300 A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: Current and emerging issues, Atmos. Chem. Phys., 9, 5155–5236, 2009.

Im, Y., Jang, M., and Beardsley, R. L.: Simulation of aromatic SOA formation using the lumping model integrated with explicit gas-phase kinetic mechanisms and aerosol-phase reactions, Atmos. Chem. Phys., 14,

305 4013–4027, 2014.

- Jenkin, M. E., Saunders, S. M., Wagner, V., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part B): Tropospheric degradation of aromatic volatile organic compounds, Atmos. Chem. Phys., 3, 181–193, 2003.
- Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen, R.,
- 310 Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global climate modelling: A review, Atmos. Chem. Phys., 5, 1053–1123, 2005.
  - Karl, T., Apel, E., Hodzic, A., Riemer, D. D., Blake, D. R., and Wiedinmyer, C.: Emissions of volatile organic compounds inferred from airborne flux measurements over a megacity, Atmos. Chem. Phys., 9, 271–285,
- **315** 2009.
  - Kroll, J. H. and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and evolution of lowvolatility organics in the atmosphere, Atmos. Environ., 42, 3593–3624, 2008.
  - Myrdal, P. B. and Yalkowsky, S. H.: Estimating pure component vapor pressures of complex organic molecules, Ind. Eng. Chem. Res., 36, 2494–2499, 1997.
- 320 Ng, N. L., Kroll, J. H., Chan, A. W. H., Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from *m*-xylene, toluene, and benzene, Atmos. Chem. Phys., 7, 3909–3922, 2007.
  - Nguyen, K. and Dabdub, D.: Two-level time-marching scheme using splines for solving the advection equation, Atmos. Environ., 35, 1627–1637, 2001.

Nguyen, K. and Dabdub, D.: Semi-Lagrangian flux scheme for the solution of the aerosol condensation/evaporation equation, Aerosol Sci. Technol., 36, 407–418, 2002.

Pankow, J. F.: An absorption model of the gas/aerosol partitioning involved in the formation of secondary organic aerosol, Atmos. Environ., 28, 189–193, 1994.

Pankow, J. F. and Asher, W. E.: SIMPOL.1: A simple group contribution method for predicting vapor pressures and enthalpies of vaporization of multifunctional organic compounds, Atmos. Chem. Phys., 8, 2773–2796,

330 2008.

325

340

- Pun, B. K., Griffin, R. J., Seigneur, C., and Seinfeld, J. H.: Secondary organic aerosol 2. Thermodynamic model for gas/particle partitioning of molecular constituents, J. Geophys. Res., 107, 4333, doi:10.1029/2001JD000542, 2002.
- Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the development of the Mas-
- 335 ter Chemical Mechanism, MCM v3 (Part A): Tropospheric degradation of non-aromatic volatile organic compounds, Atmos. Chem. Phys., 3, 161–180, 2003.

SCAQMD: Final 2012 Air Quality Management Plan, Tech. rep., South Coast Air Quality Management District, 2013.

- Seinfeld, J. and Pandis, S.: Atmospheric chemistry and physics: From air pollution to climate change, Wiley-Interscience, 2006.
- Song, C., Na, K., and Cocker, D. R.: Impact of the hydrocarbon to NO<sub>X</sub> ratio on secondary organic aerosol formation, Environ. Sci. Technol., 39, 3143–3149, 2005.

Suzuki, T., Ohtaguchi, K., and Koide, K.: Application of principal components analysis to calculate Henry's constant from molecular structure, Comput. Chem., 16, 41–52, 1992.

- 345 Wesely, M.: Parameterization of surface resistances to gaseous dry deposition in regional-scale numerical models, Atmos. Environ., 23, 1293–1304, 1989.
  - White, M., Russo, R., Zhou, Y., Ambrose, J., Haase, K., Frinak, E., Varner, R., Wingenter, O., Mao, H., Talbot, R., and Sive, B.: Are biogenic emissions a significant source of summertime atmospheric toluene in the rural Northeastern United States?, Atmos. Chem. Phys., 9, 81–92, 2009.
- 350 Xu, J., Griffin, R. J., Liu, Y., Nakao, S., and Cocker III, D. R.: Simulated impact of NO<sub>X</sub> on SOA formation from oxidation of toluene and *m*-xylene, Atmos. Environ., 101, 217–225, 2015.



**Figure 1.** Surrogate species for the new aromatic-derived SOA groups: (a) C1, (b) C2, (c) C3, (d) C4 and (e) C5.



**Figure 2.** Modeled 24-hour average total SOA concentrations (a) before updates to vapor pressure calculations and SOA lumping scheme, (b) after implementation of SIMPOL.1 for vapor pressure calculations, and (c) after adoption of SIMPOL.1 and new SOA lumping scheme. Note that the scale changes between panels.



**Figure 3.** Modeled domain-average diurnal SOA concentrations by SOA group (a) before updates to vapor pressure calculations and SOA lumping scheme, (b) after implementation of SIMPOL.1 for vapor pressure calculations, and (c) after adoption of SIMPOL.1 and new SOA lumping scheme. Note that the scale changes between panels. Surrogate structures for each of the lumped SOA species are shown in Figure 4.



**Figure 4.** Surrogate species for each SOA group after changes shown in Table 2, but before addition of new aromatic-derived species: (a) B5, (b) A5, (c) B3, (d) B2, (e) A1, (f) B1, (g) A2, (h) A3, (i) A4, (j) B7, (k) B6, and (l) B4.



**Figure 5.** Modeled 24-hr average SOA concentrations after inclusion of new aromatic chemistry assuming AROH emissions are (a) all *m*-xylene, (b) all toluene, or (c,d) 30% each *m*-xylene and toluene. Vapor pressures are calculated using (a,b,c) SIMPOL.1, or (d) SIMPOL.1 with SOA species C1 - C5 adjusted to match those used by Xu et al. (2015) at 298 K.



Figure 6. Modeled domain-average dirunal SOA concentrations by SOA group after inclusion of new aromatic chemistry assuming AROH emissions are (a) all m-xylene, (b) all toluene, or (c) 30% each m-xylene and toluene.



**Figure 7.** Modeled domain-average diurnal (a) NO and  $NO_2$  concentrations and (b) [NO] to [NO<sub>2</sub>] ratio, assuming AROH emissions are 30% each *m*-xylene and toluene.



**Figure 8.** Diurnal fraction of aromatic-derived SOA species in the particle phase, averaged over the model domain. This fraction is defined as the mass concentration of a lumped SOA species relative to its total mass concentration (gas- and particle-phase).

SOA Group	Description	Species Included <sup>1</sup>
C1	furanones	UR24, UR72 and UR75 <sup>2</sup>
C2	nitro-phenols	RPR4, RP98, UR22, UR57, UR58, UR65, UR66 and
		UR77 <sup>2</sup>
C3	peroxy-nitrates	PN11 <sup>2</sup> , PN12, PN13 and PN14
C4	epoxides	R102 <sup>2</sup> and UR76
C5	peroxy-bicyclic rings	RP29, RP30 <sup>2</sup> and RP31

Table 1. New SOA groups from addition of toluene and *m*-xylene oxidation mechanisms.

<sup>1</sup>See Xu et al. (2015) for descriptions of individual SOA species. <sup>2</sup>Surrogate species

**Table 2.** Updates to the SOA lumping scheme as part of this work, along with their motivation. The original SOA lumping scheme is described in Pun et al. (2002) and Griffin et al. (2005).

SOA Group	Change	Reason	
A1	made UR21 (ketopropanoic acid)	Species UR21 has a much higher domain average	
	the surrogate species	concentration than the previous surrogate, UR28	
		(oxalic acid)	
A2, A4	moved UR3 from A4 to A2	The calculated $p_L^O$ for UR3 is closer to species in	
		SOA group A2	
A2, A4	moved UR26 from A2 to A4	same as above	
A2	made UR3 the surrogate species	Species UR3 has representative $p_L^{\cal O}$ , Henry's law	
		constant, and structure for group A2 and has the	
		highest domain average concentration	
A4	made UR26 the surrogate species	same as above	
B4	remove UR34 from group B4	Species UR34 is no longer produced in the gas-	
		phase mechanism (Griffin et al., 2005).	
$B4, B6^1, B7^1$	split group B4 into B4 (AP12),	These three SOA species have high domain aver-	
	B6 (AP11), and B7 (UR20)	age concentrations and calculated $p_L^O$ (298 K) that	
		vary by an order of magnitude or more.	

<sup>1</sup>New lumped SOA species

	Original			Original	1
SOA	CIT	aroCACM/	SOA	СІТ	aroCACM/
Group	Model	MPMPO	Group	Model	ΜΡΜΡΟ
A1	UR21	UR21*	B4	AP11	AP12*
	UR28*	UR28		AP12*	
				UR20	
A2	RP13	RP13		UR34	
	RP17	RP17			
	RP18*	RP18	B5	AP8	AP8
	UR29	UR29		UR5	UR5
	UR30	UR30		UR6	UR6
	UR26	UR3*		AP7*	AP7*
A3	RPR9*	RPR9*	B6		AP11*
-	RP12	RP12	-		
			B7		UR20*
A4	UR3*	UR26*			
	UR23	UR23	C1		UR75*
	UR8	UR8			UR24
					UR72
A5	UR17*	UR17*			
	UR7	UR7	C2		UR77*
					RP98
B1	AP1	AP1			RPR4
	AP6*	AP6*			UR22
	UR31	UR31			UR66
					UR57
B2	ADAC	ADAC			UR58
	RPR7*	RPR7*			UR65
	RP14	RP14			
	RP19	RP19	C3		PN11*
	UR2	UR2			PN12
	UR14	UR14			PN13
	UR27	UR27			PN14
В3	AP10*	AP10*	C4		R102*
	UR11	UR11			UR76
	UR15	UR15			
			C5		RP30*
					RP29
					RP31

Table S1. Component species and surrogate structures of SOA groups in the original and updated CIT model. Surrogate species are indicated with a (\*).

	Average (μg/m3)	1 Standard Deviation (μg/m³)
C1	2.36E-01	4.24E-01
C2	3.54E-02	6.18E-02
C3	1.03E-01	1.30E-01
C4	9.42E-03	1.88E-02
C5	1.55E-02	2.00E-02

Table S2. Domain and 24 hour average concentrations of aromatic-derived lumped SOA species C1 – C5 in the 30% each m-xylene and toluene scenario.