1 The format of this reply to Referee 1 is as follows: 2 In the first part, "Authors' Response to Anonymous Referee 1", we 3 4 provide a point-by-point response to the referee's comments. We provide each of the 5 referee's comments in **bold font**. After each comment that requires a response we provide a response (in regular font). If we modified the manuscript in response to a 6 7 comment, we describe what the modification was, and indicate where it was made in the 8 revised manuscript (the revised manuscript is provided at the end of this document). 9 In the second section, Additional modifications to the manuscript, we 10 11 describe modifications to the manuscript not made in response to any specific comment 12 of either referee. For the most part these modifications are minor, however we did fix 13 two errors in the final analytical solution (errors in the text, not the code; so these did 14 not affect any of the presented results). 15

In the third section, we provide the revised manuscript.

16

Authors' Response to Anonymous Referee 1

The authors derive an analytical solution of the two-dimensional steady Boussinesq equations in the limit of zero Reynolds number. To obtain this solution they transform the equations into a sixth-order equation for the streamfunction. Then they seek a solution in the form of a single-harmonic, $A(z)\cos(kx)$. Then the general solution is found as an infinite summation of single-harmonic solutions. The procedure is rigorous and well-described. In the second part of the paper the authors proceed to use this analytical solutions to test different boundary conditions for the pressure Poisson equation in a particular numerical implementation. They consider two cases in which they compare numerical results for homogeneous and inhomogeneous types boundary conditions for the pressure Poisson equation to the analytical solution. Only the inhomogeneous boundary condition passes both tests. In summary, this is an interesting paper that deserves publication. I have several minor comments:

1. Section 2.1: Please explain the physical meaning of equation (3), i.e. that the equation for b is based on the transport equation for the temperature.

Equation (2.3) is the thermal energy equation (differential form of the first law of thermodynamics). In the revised manuscript we now briefly discuss the governing equations, including (2.3) in the paragraph right after (2.4). We also give a reference to where the equations are described more fully (Chandrasekhar 1961). Further down in the paragraph we give a reference (Kundu 1990) for the Brunt-Väisälä frequency N, a parameter that appears in (2.3).

2. Please specify which condition has been substituted in which equation in order to obtain Eqs. (24-25).

Right before (2.24) we now state that: "In view of (2.7) and (2.17), the impermeability condition w(x,0) = 0 and no-slip condition u(x,0) = 0 yield"

3. Section 2.3: "The derivation of the u field requires considerable effort and is not pursued." I do not understand this. As far as I understand, the analytical solution of u is

simply the analytical partial derivative of Eq. (39) with respect to x, which can be explicitly written as an infinite sum. Do I miss anything?

The referee is not quite right. As can be seen from (7), one gets u by differentiating the streamfunction (specified in (39)) with respect to z, not x [One gets w by differentiating the streamfunction with respect to x, and that's a relatively easy calculation]. Moreover, as can be seen from (31), z is implicitly inside four of the factors in (39). This is why obtaining u is potentially complicated.

However, we decided to go ahead and take the z-derivative of (39) and see if we could obtain a relatively compact final form for u. It turns out that by making use of addition formulas for sines, the results do simplify considerably and the final form of u is not too bad. So, we now present the analytical solution for u: (2.32) for the single-harmonic wave and (2.41) for the square wave. The differences between the analytical u and the u obtained by finite differencing the streamfunction are visually imperceptible for tests A-1 and A-2. The quantitative changes are generally less than 1 % but are on the order of 1 % near the surface. Our reported values for the R ratios do not change for A-1 but there is a minor change in one of the R ratios for A-2: R_b was originally $R_b \cong 3.6 \times 10^{-3}$ but is now $R_b \cong 3.8 \times 10^{-3}$. The refined value is given in the revised manuscript. With the analytical solution for u now provided, there is now no reason to discuss obtaining the u field by finite-differencing the streamfunction. We have omitted such prose from the revised manuscript.

4. Section 3: "The surface condition on pressure is the inhomogeneous Neumann condition that arises from projecting the vertical equation of motion into the vertical, and imposing the impermeability condition (Vreman, 2014; also see our Appendix)." The sentence can be maintained, but a sentence should be added that, in addition, it is important that the discretized Poisson equation somehow incorporates the condition that delta = div u=0 on the wall or in the direct vicinity of the wall. This was also stressed by Vreman (and others) and is briefly mentioned in the appendix. It is good to include this requirement also in the main text. In the method of the authors the condition delta=0 near the wall is probably implicitly enforced via the alternative Poisson equation, specified in the Appendix, Eq. (A3b). After Eq. (A3b) the authors cite the pressure Poisson equation parodox using a sentence of Gresho and Sani. Please mention there that Vreman has revisited this paradox

and has shown that, at least for the standard staggered method, the discretized version of (A3b) (with appropriate Neumann condition) is equivalent to the discretized version of (A3a) supplemented with the condition that div(Laplacian(u))=0 in the direct vicinity of the wall. Equipped with the latter the condition, the diffusion equation d delta/dt = nu * Laplacian(delta) leads to delta=0 for all time.

Yes, the divergence-free condition near the wall is enforced via the alternative Poisson equation (A3b). We have modified the sentence right after (A3b) to emphasize that this alternative Poisson equation assures that (A2) (the divergence-free condition) is satisfied.

Since we want section 3 to focus on the verification tests, we did not want to divert too much attention in that section to the technical details of the pressure equation, pressure boundary condition, and related topics. However, we agree that these details are certainly important and should be clearly discussed. So, as a compromise, we modified a sentence in section 3 to read: "The pressure is diagnosed from a Poisson equation (equation (3b), discussed in the Appendix),..." Then, in the Appendix we modified the sentence about (A3b) enforcing the divergence-free condition (the modification described in the paragraph above) and added two sentences to the discussion of (A3b) along the lines suggested by the referee. Then, after those two new sentences, we end the paragraph with: "We note that (A3b) is the form adopted in our numerical code."

5. Section 3: The numerical solutions are obtained on an un-staggered grid. Please explain what was done to prevent odd-even decoupling of the pressure. Was the Rhie-Chow interpolation method used, for example?

The numerical solutions are obtained on a staggered (Arakawa C) grid so no decoupling of the pressure was occurring/noticed. We now mention "staggered (Arakawa C) grid" in the discussion of the DNS code in section 3.

6. Section 3: Please explain the meaning of the abbreviations HNC and INC (I guess homogeneous Neumann condition and inhomogeneous Neumann condition).

Yes, HNC is out acronym for homogeneous Neumann condition and INC is our acronym for inhomogeneous Neumann condition. We now introduce these acronyms near the end of the second paragraph of section 3.

110	Additional modifications to the manuscript (i.e., not made in
111	response to the reviewers' comments)
112	
113	Please note that the third author would like his middle initial "A" included in his name:
114	Jeremy A. Gibbs.
115	
116	We have added a new reference: Egger (1981). The Egger study was related to ours in
117	that it was concerned with a linear analysis of the 2D Boussinesq governing equations
118	for thermally driven flow. Egger's analysis was largely for slope flows, though with flat
119	terrain (our focus) considered as a special case. However, Egger outlines how to get the
120	analytical solution but does not actually provide the final analytical solution. We
121	mention this Egger study in the second paragraph of Section 1. We also mention it in
122	the paragraph right after (2.8): the restriction on acceptable surface buoyancies
123	described in that paragraph was first noted by Egger, though without details.
124	
125	A correction was made to the original equations (2.39) and (2.40) [these now appear as
126	equations (2.40) and (2.42), respectively]. The factor $k^{1/3}$ in the denominator of the
127	term in front of the summation in (2.39) and the factor $k^{2/3}$ in the numerator of the
128	term in front of the summation in (2.40) should be kept inside the summations. These
129	factors were treated correctly in the computer code, so none of the presented results
130	were affected.
131	
132	Section 3. We now make the number of points in the x and z direction unambiguous:
133	instead of writing the number of points in test A-1 as (513, 1025) we write, "consisted
134	of 513 points in the x direction and 1025 points in the z direction," Similarly, for test
135	A-2, we now write, "was generated with 2049 points in the x direction and 513 points

in the z direction,..." In the Appendix we now write the time step as Δt instead of δt since the symbol δ has already been used to represent the divergence of the velocity field. In several places in the manuscript we now use bold to indicate the vector \mathbf{u} (formerly we used \vec{u}). We have slightly modified the acknowledgements statement (we now thank the anonymous reviewer).

An analytical verification test for numerically simulated convective flow above a thermally heterogeneous surface

by Alan Shapiro, Evgeni Fedorovich, and Jeremy A. Gibbs

Abstract. An analytical solution of the Boussinesq equations for the motion of a viscous stably stratified fluid driven by a surface thermal forcing with large horizontal gradients (step changes) is obtained. This analytical solution is one of the few available for wall-bounded buoyancy-driven flows. The solution can be used to verify that computer codes for Boussinesq fluid system simulations are free of errors in formulation of wall boundary conditions and to evaluate the relative performances of competing numerical algorithms. Because the solution pertains to flows driven by a surface thermal forcing, one of its main applications may be for testing the no-slip, impermeable wall boundary conditions for the pressure Poisson equation. Examples of such tests are presented.

1 Introduction

161

162

163

164

165

166

167

168

169

170

171

172

173

174

175

176

177

178

179

180

Thermal disturbances associated with variations in underlying surface properties can drive local circulations in the atmospheric boundary layer (Atkinson, 1981; Briggs, 1988; Hadfield et al., 1991; Segal and Arritt, 1992; Simpson, 1994; Mahrt et al., 1994; Pielke, 2001; McPherson, 2007; Kang et al., 2012) and affect the development of the convective boundary layer (Patton et al., 2005; van Heerwaarden et al., 2014). Computational fluid dynamics (CFD) codes for modeling such flows commonly solve the Boussinesq equations of motion and thermal energy for a viscous/diffusive stably stratified fluid. In this paper we present an analytical solution of the Boussinesq equations for flows driven by a surface thermal forcing with large gradients (step changes) in the horizontal. The solution can be used to verify that CFD codes for Boussinesq fluid system simulations are free of errors, and to evaluate the relative performances of competing numerical algorithms. Such verification procedures are important in the development of CFD models designed for research, operational, and classroom applications.

We solve the linearized Navier-Stokes and thermal energy equations analytically for the case where the surface buoyancy varies laterally as a square wave (Fig. 1). Attention is restricted to the steady state. No boundary-layer approximations are made; the solution is non-hydrostatic, and both horizontal and vertical derivatives are included in the viscous stress and thermal diffusion terms. The solution is similar to that of Axelsen et al. (2010) for katabatic flow above a cold strip, but is easier to evaluate (no

slope present) and applies to the more general scenario where the viscosity and diffusivity coefficients can differ. The flow is also similar to a special case (no slope) considered by Egger (1981), although a final analytical solution was not provided in that study. Strictly speaking, the linearized Navier-Stokes equations apply to a class of very low Reynolds number motions known as creeping flows. Such flows appear in studies of lubrication, locomotion of microorganisms, lava flow, and flow in porous media. Of course, for the task at hand, if our linear solution is to serve as a benchmark for a nonlinear numerical model solution, it is essential that the parameter space be restricted to values for which the model's nonlinear terms are negligible.

Because the solution pertains to flows driven by a surface thermal forcing, one of its main applications may be as a test for surface boundary conditions in the pressure Poisson equation. In models of atmospheric boundary layer flows, the buoyancy is a major contributor to the forcing term in the Poisson equation and also appears in the associated surface boundary condition. The pressure boundary condition on a solid boundary in incompressible (Boussinesq) fluid flows is an important and complex issue that has long been fraught with technical difficulties and controversies (Strikwerda, 1984; Orszag et al., 1986; Gresho and Sani, 1987; Gresho, 1990; Temam, 1991; Henshaw, 1994; Petersson, 2001; Sani et al., 2006; Rempfer, 2006; Guermond et al., 2006; Nordström et al., 2007; Shirokoff and Rosales, 2011; Hosseini and Feng, 2011; Vreman, 2014). Typical fractional-step solution methodologies and associated pressure (or

pseudo-pressure) boundary-condition implementations are often verified using various prototypic flows such as Poiseuille flows, lid-driven cavity flows, flows over cylinders or bluff bodies, viscously decaying vortices, and dam-break flows. We are unaware of verification tests in which flows were driven by a heterogeneous surface buoyancy forcing. Our solution is designed to fill this gap.

The analytical solution is derived in Sect. 2. In Sect. 3, this solution is compared to numerically simulated fields in a steady state. Two versions of a numerical code are run: a version in which the correct surface pressure boundary condition is applied, and a version in which the pressure condition is mis-specified. A summary follows in Sect. 4.

2 Analytical solution

We derive the solution for steady flow over an underlying surface along which the buoyancy varies laterally as a single harmonic function. This single-harmonic solution is then used as a building block in a Fourier representation of the square-wave solution.

2.1 Governing equations

Consider the flow of a viscous stably stratified fluid that fills the semi-infinite domain above a solid horizontal surface (placed at z=0). This surface undergoes a steady thermal forcing that varies periodically in the right-hand Cartesian x direction, but is independent of the y direction. The two-dimensional (x, z) flow is periodic in x, and

satisfies the linearized (assuming the disturbance is of small amplitude) governing equations under the Boussinesq approximation,

$$0 = -\frac{\partial \Pi}{\partial x} + \nu \nabla^2 u \,, \tag{2.1}$$

$$224 0 = -\frac{\partial \Pi}{\partial z} + b + \nu \nabla^2 w, (2.2)$$

$$0 = -N^2 w + \alpha \nabla^2 b , \qquad (2.3)$$

$$\frac{\partial u}{\partial x} + \frac{\partial w}{\partial z} = 0. \tag{2.4}$$

227 Apart from notational differences, (2.1)–(2.4) are the two-dimensional steady state 228 versions of (55)–(57) of Sect. II of Chandrasekhar (1961). Equations (2.1) and (2.2) are 229 the horizontal (x) and vertical (z) equations of motion, respectively, (2.3) is the thermal 230 energy equation (differential form of the first law of thermodynamics) expressed in 231 terms of the buoyancy variable (defined below), and (2.4) is the incompressibility 232 condition. Here u and w are the horizontal and vertical velocity components, $\Pi \equiv$ $[p-p_e(z)]/\rho_w$ is the kinematic pressure perturbation [p is pressure, $p_e(z)$ is pressure in a 233 hydrostatic environmental state in which the density profile is $\, \rho_e(z) \, , \, \, \rho_w \,$ is a constant 234 reference density, say, $\rho_e(0)$], and $b \equiv -g[\rho - \rho_e(z)]/\rho_w$ is the buoyancy, where ρ is the 235 236 actual density, and g is the acceleration due to gravity. The Brunt-Väisälä frequency $N \equiv \sqrt{-(g/\rho_w)d\rho_e/dz}$ of the ambient fluid (Kundu 1990), kinematic viscosity $\nu\,,$ and 237 238 thermal diffusivity α are taken constant.

We obtain our solution using a standard vorticity/streamfunction formulation.

240 Cross-differentiating (2.1) and (2.2) yields the vorticity equation,

$$0 = -\frac{\partial b}{\partial x} + \nu \nabla^2 \eta \,, \tag{2.5}$$

242 where $\eta \equiv \partial u/\partial z - \partial w/\partial x$ is the vorticity. Eliminating b from (2.3) and (2.5) yields

$$\nabla^4 \eta = \frac{N^2}{\nu \alpha} \frac{\partial w}{\partial x}.$$
 (2.6)

244 Introducing a streamfunction ψ defined through

245
$$u = \partial \psi / \partial z, \qquad w = -\partial \psi / \partial x,$$
 (2.7)

guarantees that (2.4) is satisfied, and transforms (2.6) into a single equation for ψ ,

$$\nabla^6 \psi + \frac{N^2}{\nu \alpha} \frac{\partial^2 \psi}{\partial x^2} = 0. \tag{2.8}$$

The dependent variables are assumed to vanish far above the surface $(z \to \infty)$. On the 248 249 surface we apply no-slip (u=0) and impermeability (w=0) conditions, and specify a 250 periodic (in x) buoyancy distribution. As we will now see, restricting the dependent variables to steady periodic forms that vanish as $z \to \infty$ also restricts acceptable 251 252 distributions of the surface buoyancy. The restriction was first noted by Egger (1981, 253 Sect. 3c), though without details. Averaging (2.3) over one period (using $w = -\partial \psi/\partial x$) yields $d^2\overline{b}/dz^2=0$, which integrates to $\overline{b}=A+Bz$ (\overline{b} is the average of b; A and B are 254 constants). Taking $b\to 0$ as $z\to \infty$, implies that $\overline{b}\to 0$ as $z\to \infty$, in which case A=255 B=0, and $\overline{b}(z)=0$. In particular, at the surface, $\overline{b}(0)=0$. If a surface distribution 256

b(x,0) violates this condition, the ground acts as a net heat source/sink. In an unsteady model, such a source/sink would force a continually upward-developing disturbance, and a steady state could never be attained.

260

261

2.2 Single-harmonic forcing

For a surface buoyancy of the form $b(x,0) \propto \sin kx$, (2.3) indicates that ψ is of the form

$$\psi = A(z)\cos kx \,. \tag{2.9}$$

264 Application of (2.9) in (2.8) yields

$$\left(\frac{d^2}{dz^2} - k^2\right)^3 A - \frac{N^2 k^2}{\nu \alpha} A = 0,$$
 (2.10)

266 which has solutions of the form $A \propto e^{Mz}$ for M satisfying

$$(M^2 - k^2)^3 = \frac{N^2 k^2}{\nu \alpha}. \tag{2.11}$$

268 Taking the one-third power of (2.11) yields a useful intermediate result:

269
$$M^2 - k^2 = \frac{N^{2/3} k^{2/3}}{\nu^{1/3} \alpha^{1/3}} e^{2n\pi i/3}, \qquad (2.12)$$

270 where n is an integer. Rearranging (2.12) and taking the square root yields

271
$$M = \pm \sqrt{k^2 + \frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}}} e^{2n\pi i/3} . \tag{2.13}$$

272 Equation (2.13) furnishes six roots, two for each of n = 0, 1, 2. To ensure that $A(z) \rightarrow 0$

273 as $z \to \infty$, we reject the roots with a positive real part. With the radicand of (2.13)

274 expressed in polar form, the physically acceptable roots are

275
$$M_0 = -\sqrt{k^2 + \frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}}}, \quad (n = 0),$$
 (2.14a)

276
$$M_1 = -r^{1/2}e^{i\phi/2}, \qquad (n=1),$$
 (2.14b)

$$277 \hspace{1cm} M_2 = -r^{1/2} e^{-i\phi/2} \,, \hspace{1cm} (n=2), \hspace{1cm} (2.14c)$$

278 where the subscript on M denotes the associated value of n, and r and ϕ are defined by

$$r \equiv \sqrt{\left[k^2 + \frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}}\cos\left(\frac{2\pi}{3}\right)\right]^2 + \left[\frac{N^{2/3}k^{2/3}}{\nu^{1/3}\alpha^{1/3}}\sin\left(\frac{2\pi}{3}\right)\right]^2} ,$$
 (2.15)

$$\cos \phi = \frac{1}{r} \left[k^2 + \frac{N^{2/3} k^{2/3}}{\nu^{1/3} \alpha^{1/3}} \cos \left(\frac{2\pi}{3} \right) \right], \qquad \sin \phi = \frac{1}{r} \left(\frac{N^{2/3} k^{2/3}}{\nu^{1/3} \alpha^{1/3}} \right) \sin \left(\frac{2\pi}{3} \right) > 0. \tag{2.16}$$

- While solving (2.16) for ϕ , care must be taken when evaluating arcsin or arccos 282 functions that ϕ appears in the correct quadrant (ϕ should be in quadrant I or II so
- $\phi/2$ should always be in quadrant I). Also note from (2.14b) and (2.14c) that $\,M_2^{}\,$ is the 283
- complex conjugate of $\,M_1\,\,(\,M_2=M_1^{\,*}),$ a fact that will often be used below. 284
- 285 With the general solution for ψ written as

281

286
$$\psi = (Be^{M_0z} + Ce^{M_1z} + De^{M_2z})\cos kx, \qquad (2.17)$$

287 where B, C, and D are constants, the vorticity becomes,

288
$$\eta = \left[B(M_0^2 - k^2) e^{M_0 z} + C(M_1^2 - k^2) e^{M_1 z} + D(M_2^2 - k^2) e^{M_2 z} \right] \cos kx, \qquad (2.18)$$

289 and the buoyancy follows from (2.3) as

290
$$b = \frac{kN^2}{\alpha} \left(\frac{B}{M_0^2 - k^2} e^{M_0 z} + \frac{C}{M_1^2 - k^2} e^{M_1 z} + \frac{D}{M_2^2 - k^2} e^{M_2 z} \right) \sin kx + b_h, \qquad (2.19)$$

291 where $\nabla^2 b_h = 0$. In view of (2.12), equation (2.19) becomes

292
$$b = \frac{k^{1/3} \nu^{1/3} N^{4/3}}{\sigma^{2/3}} (B e^{M_0 z} + e^{-2\pi i/3} C e^{M_1 z} + e^{-4\pi i/3} D e^{M_2 z}) \sin kx + b_h.$$
 (2.20)

- 293 Applying (2.18) and (2.20) in (2.5) yields an equation for $\partial b_h/\partial x$, which upon use of
- 294 (2.12) and $M_2=M_1^{\ *}$ reduces to $\partial b_h/\partial x=0$. So b_h is, at most, a function of z. Since
- 295 $\nabla^2 b_h = 0$, b_h is, at most, a linear function of z, and since b should vanish as $z \to \infty$,
- 296 that linear function must be 0. Thus, $b_h = 0$.
- The pressure follows from (2.1) and (2.12) as

$$\Pi = \frac{\nu^{2/3} N^{2/3}}{k^{1/3} \alpha^{1/3}} \left(B M_0 \, e^{M_0 z} + C M_1 e^{2\pi i/3} e^{M_1 z} + D M_2 e^{4\pi i/3} e^{M_2 z} \right) \sin kx + G(z) \,, \qquad (2.21)$$

- where G(z) is a function of integration. Applying (2.21) in (2.2), and using (2.11) yields
- 300 dG/dz = 0, so G is constant. For Π to vanish as $z \to \infty$, this constant must be zero.
- 301 The surface conditions determine B, C, and D. The surface buoyancy is

302
$$b(x,0) = b_0 \sin kx, \qquad (2.22)$$

303 where b_0 is a constant forcing amplitude. Application of (2.20) in (2.22) yields

304
$$B + e^{-2\pi i/3} C + e^{-4\pi i/3} D = \frac{b_0 \alpha^{2/3}}{k^{1/3} \nu^{1/3} N^{4/3}}.$$
 (2.23)

305 In view of (2.7) and (2.17), the impermeability condition w(x,0) = 0 and no-slip

306 condition u(x,0) = 0 yield

$$307 B+C+D=0, (2.24)$$

$$BM_0 + CM_1 + DM_2 = 0. (2.25)$$

309 Straightforward but lengthy manipulations yield the solution of (2.23)–(2.25):

$$B = -\left(\frac{b_0 \, \alpha^{2/3}}{\sqrt{3} \, k^{1/3} \nu^{1/3} N^{4/3}}\right) \frac{2 r^{1/2} \mathrm{sin}(\phi/2)}{M_0 + 2 r^{1/2} \mathrm{cos}(\pi/3 + \phi/2)}\,, \tag{2.26}$$

$$C = -i \left[\frac{b_0 \, \alpha^{2/3}}{\sqrt{3} \, k^{1/3} \nu^{1/3} N^{4/3}} \right] \frac{M_2 - M_0}{M_0 + 2 r^{1/2} \text{cos}(\pi/3 + \phi/2)}, \tag{2.27}$$

$$D = i \left(\frac{b_0 \, \alpha^{2/3}}{\sqrt{3} \, k^{1/3} \nu^{1/3} N^{4/3}} \right) \frac{M_1 - M_0}{M_0 + 2 r^{1/2} \mathrm{cos}(\pi/3 + \phi/2)} \,. \tag{2.28}$$

313 Applying (2.26)–(2.28) in (2.17), (2.20), and (2.18), with (2.12) used in the latter

314 equation, and noting that B is real, while $D=C^*$ (since $M_2=M_1^*$), we obtain

315
$$b = \frac{2b_0}{\sqrt{3}} \frac{e^{-Z_c} \left[\mu \cos(Z_s + \pi/6) + \cos(Z_s + \pi/6 + \phi/2)\right] - e^{M_0 z} \sin(\phi/2)}{\mu + 2\cos(\pi/3 + \phi/2)} \sin kx, \qquad (2.29)$$

316
$$\psi = \frac{2b_0 \alpha^{2/3}}{\sqrt{3} k^{1/3} \nu^{1/3} N^{4/3}} \frac{e^{-Z_c} [\mu \sin Z_s + \sin(Z_s + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2\cos(\pi/3 + \phi/2)} \cos kx, \qquad (2.30)$$

317 where

Application of (2.30) in (2.7) yields the velocity components as

320
$$u = \frac{2b_0 \alpha^{2/3} r^{1/2}}{\sqrt{3} k^{1/3} \nu^{1/3} N^{4/3}} \frac{e^{-Z_c} [\mu \sin(\phi/2 - Z_s) - \sin Z_s] - \mu e^{M_0 z} \sin(\phi/2)}{\mu + 2\cos(\pi/3 + \phi/2)} \cos kx$$
 (2.32)

321
$$w = \frac{2b_0 \alpha^{2/3} k^{2/3}}{\sqrt{3} \nu^{1/3} N^{4/3}} \frac{e^{-Z_c} [\mu \sin Z_s + \sin(Z_s + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2\cos(\pi/3 + \phi/2)} \sin kx.$$
 (2.33)

323 2.3 Piecewise constant (square wave) forcing

- 324 Next, consider the case where the surface buoyancy varies horizontally as a square
- 325 wave, with a distribution over one period L given by

326
$$b(x,0) = \begin{cases} b_{\text{max}}, & 0 < x < L/2, \\ -b_{\text{max}}, & L/2 < x < L. \end{cases}$$
 (2.34)

327 Such a distribution can be expressed as the Fourier series:

328
$$b(x,0) = \sum_{n=1}^{\infty} b_n \sin\left(\frac{n\pi x}{L}\right), \tag{2.35}$$

$$b_n = \frac{2}{L} \int_0^L b(x,0) \sin\left(\frac{n\pi x}{L}\right). \tag{2.36}$$

330 Application of (2.34) in (2.36) yields

331
$$b_n = \frac{2b_{\text{max}}}{n\pi} \left[1 - 2\cos(n\pi/2) + \cos(n\pi) \right]. \tag{2.37}$$

- 332 The solutions for b, ψ, u , and w can then be written as summations over the single-
- 333 harmonic solutions (2.29), (2.30), (2.32), and (2.33), with k related to n by

$$334 k = \frac{n\pi}{L}, (2.38)$$

335 and with b_0 replaced by b_n :

336
$$b = \frac{2}{\sqrt{3}} \sum_{n=1}^{\infty} b_n \frac{e^{-Z_c} \left[\mu \cos(Z_s + \pi/6) + \cos(Z_s + \pi/6 + \phi/2)\right] - e^{M_0 z} \sin(\phi/2)}{\mu + 2\cos(\pi/3 + \phi/2)} \sin\left(\frac{n\pi x}{L}\right), \quad (2.39)$$

337
$$\psi = \frac{2\alpha^{2/3}}{\sqrt{3}\nu^{1/3}N^{4/3}} \sum_{n=1}^{\infty} \frac{b_n}{k^{1/3}} \frac{e^{-Z_c} [\mu \sin Z_s + \sin(Z_s + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2\cos(\pi/3 + \phi/2)} \cos\left(\frac{n\pi x}{L}\right), \quad (2.40)$$

338
$$u = \frac{2\alpha^{2/3}}{\sqrt{3}\nu^{1/3}N^{4/3}} \sum_{n=1}^{\infty} b_n \frac{r^{1/2}}{k^{1/3}} \frac{e^{-Z_c} [\mu \sin(\phi/2 - Z_s) - \sin Z_s] - \mu e^{M_0 z} \sin(\phi/2)}{\mu + 2\cos(\pi/3 + \phi/2)} \cos\left(\frac{n\pi x}{L}\right), (2.41)$$

339
$$w = \frac{2\alpha^{2/3}}{\sqrt{3}\nu^{1/3}N^{4/3}} \sum_{n=1}^{\infty} b_n k^{2/3} \frac{e^{-Z_c} [\mu \sin Z_s + \sin(Z_s + \phi/2)] - e^{M_0 z} \sin(\phi/2)}{\mu + 2\cos(\pi/3 + \phi/2)} \sin\left(\frac{n\pi x}{L}\right).$$
 (2.42)

341 3 Verification tests

A solution of the linearized equations may be used to verify a nonlinear code if the nonlinear terms are sufficiently small. Unfortunately, a priori estimates of such terms expressed, for example, through a Reynolds number, are not straightforward since the relevant velocity and length scales in our problem are only evident after a solution has been obtained. We thus seek an appropriate set of test parameters through trial and error, guided by a posteriori linear solution estimates of the terms $\mathbf{u} \cdot \nabla b$ and $\mathbf{u} \cdot \nabla \eta$ $[\mathbf{u} = (u, w)]$ present in nonlinear versions of (2.3) and (2.5), respectively. Specifically, for any computed candidate solution, we formed the ratios of the largest values of those nonlinear terms to the largest values of the corresponding linear terms, that is, the

terms actually present in (2.3) and (2.5). We need only consider one such linear term

per ratio since (2.3) and (2.5) are comprised of two terms of equal magnitude. A

solution was deemed to be sufficiently linear if

354
$$R_{\eta} \equiv \frac{\max \left| \mathbf{u} \cdot \nabla \eta \right|}{\max \left| \partial b / \partial x \right|} < \varepsilon, \quad \text{and} \quad R_{b} \equiv \frac{\max \left| \mathbf{u} \cdot \nabla b \right|}{\max \left| \alpha \nabla^{2} b \right|} < \varepsilon , \tag{3.1}$$

where ε (<< 1) is a prescribed threshold. The suitability of this approach was confirmed by the very close agreement between the analytical solutions and the numerical solutions obtained with the correct surface pressure condition.

The numerical model employed in our tests is a variant of a direct numerical simulation (DNS) code used in the boundary-layer and slope-flow studies of Fedorovich et al. (2001), Fedorovich and Shapiro (2009a,b), and Shapiro and Fedorovich (2013, 2014). The model solves the Boussinesq governing equations on a staggered (Arakawa C) grid. Although designed for three-dimensional simulations, the model was run in a two-dimensional (x, z) mode. The overall solution procedure is patterned on a fractional step method proposed by Chorin (1968). In our version, the prognostic equations are integrated using a filtered leapfrog scheme with explicit treatment of the viscous term. The pressure is diagnosed from a Poisson equation (equation (A3b), discussed in the Appendix), which is solved using a fast Fourier transform technique in horizontal planes, and a tridiagonal matrix inversion in the vertical. The surface condition on pressure is the inhomogeneous Neumann condition (INC) that arises from projecting the

vertical equation of motion into the vertical, and imposing the impermeability condition (Vreman, 2014; also see the Appendix). We also run a version of the code in which the surface pressure condition is mis-specified as a homogeneous Neumann condition (HNC). We hasten to add, however, that our implementation of the HNC may be quite different from implementations described in the literature. We elaborate on these technical differences and review general aspects of the problem of surface pressure specification in the Appendix.

The analytical solution was evaluated on an un-staggered (x, z) grid extending over one period of the square wave (x = 0 to x = L). The series were truncated at 50000 terms. The governing parameters were adjusted so that the linearity criteria were satisfied in comparisons with $\varepsilon = 5 \times 10^{-3}$.

In the first test, we set $\nu = \alpha = 0.001 \,\mathrm{m^2 \, s^{-1}}$, $N = 0.02 \,\mathrm{s^{-1}}$, $L = 5.12 \,\mathrm{m}$, and b_{max} $=1\times10^{-5}\,\mathrm{m\,s^{-2}}$. For the analytical solution A-1, the $(x,\,z)$ grid consisted of 513 points in the x direction and 1025 points in the z direction, with grid spacings $\Delta x = \Delta z = 0.01\,\mathrm{m}$. The linearity criteria (3.1) were satisfied with $R_\eta \cong 8.2 \times 10^{-5}$ and $R_b \cong 2.8 \times 10^{-3}$. The analytical b and w fields shown in Fig. 2 depict a broad zone of ascent above the warm surface and a compensating zone of descent over the cold surface, roughly for $z < 1.8 \,\mathrm{m}$. In the upper part of these zones (at roughly $0.9\,\mathrm{m} < z < 1.8\,\mathrm{m}$), adiabatic expansion/compression has reversed the senses of the

buoyancy fields. Surprisingly, the numerical fields in the inhomogeneous INC-1 and homogeneous HNC-1 cases are very similar to each other and to the A-1 fields. The *u* fields from A-1, INC-1, and HNC-1 shown in Fig. 3 are visually indistinguishable from one another.

To understand why the INC-1 and HNC-1 simulations are so similar, and to identify simulation parameters that might evince more substantial differences, we consider the idealized problem in which a specified buoyancy $b = b_0 e^{-\gamma z} \sin kx$ ($\gamma = h^{-1}$, where h is the e-folding depth scale) is the only forcing term in the Poisson equation $\nabla^2 \Pi = \partial b/\partial z$, with Neumann surface condition $\partial \Pi/\partial z|_0 = b(x,0)$. This idealized problem is solved as

399
$$\Pi_{\text{INC}}^* = \frac{b_0}{\gamma^2 - k^2} \left(k e^{-kz} - \gamma e^{-\gamma z} \right) \sin kx . \tag{3.2}$$

400 The corresponding solution obtained with the homogeneous Neumann condition, $\partial \Pi/\partial z \Big|_0 = 0 \,, \, {\rm is}$

402
$$\Pi_{\text{HNC}}^* = \frac{b_0}{\gamma^2 - k^2} \left(\frac{\gamma^2}{k} e^{-kz} - \gamma e^{-\gamma z} \right) \sin kx \,. \tag{3.3}$$

The relative error (RE) in the vertical pressure gradient force associated with (3.2) and (3.3), defined as the local absolute error in that force divided by the local buoyancy, is calculated as

406
$$RE \equiv \left| \frac{\partial \Pi_{\text{INC}}^* / \partial z - \partial \Pi_{\text{HNC}}^* / \partial z}{b} \right| = e^{(a-1)kz}, \qquad (3.4)$$

where $a \equiv \gamma/k$. Written in terms of the depth scale h and wavelength $\lambda = 2\pi/k$, a can be interpreted as an aspect ratio characterizing the width to depth scales of the disturbance, $a = \lambda/(2\pi h) \propto \lambda \gamma$. From (3.4) we see that RE decreases exponentially with z for disturbances characterized by small aspect ratios, a < 1 (which we refer to as deep disturbances) and increases exponentially with z for disturbances characterized by large aspect ratios, a > 1 (which we refer to as shallow disturbances). The buoyancy in Fig. 2 is suggestive of a < 1, which indicates that the first test could be classified as a deep (error-forgiving) simulation.

The preceding analysis suggests that simulations with shallow thermal disturbances (a > 1) might yield large differences between cases with inhomogeneous and homogeneous Neumann conditions. There did not appear to be a straightforward way to increase the effective a by systematically varying the parameters (e.g., increasing L tended to increase the effective h), but a set of suitable parameters were identified through trial and error and were used as the basis for the second test case.

In the second test, we set $\nu=\alpha=0.0001\,\mathrm{m^2s^{-1}},\ N=0.2\,\mathrm{s^{-1}},\ L=10.24\,\mathrm{m}$, and $b_{\mathrm{max}}=5\times10^{-6}\,\mathrm{ms^{-2}}.$ The analytical solution A-2 was generated with 2049 points in the x direction and 513 points in the x direction, with grid spacings of x0.

424 The linearity criteria were satisfied with $R_{\eta}\cong 4.8\times 10^{-5}$ and $R_b\cong 3.8\times 10^{-3}$. In

contrast to the counter-rotating convection rolls seen in the first test, the analytical b and w fields shown in Fig. 4 depict narrow updraft/downdraft pairs straddling the buoyancy discontinuities. Between the narrow updrafts is a broad region of relatively weak ascent. The w and b fields above the cold surface are mirror images of the fields above the warm surface. Note the change in the scales of the x and (especially) the z axes between Figs. 4 and 2: the low-level thermal disturbance in the second test is much shallower than the disturbance in the first test (and is suggestive of a > 1). In this second test case we find dramatic differences between the inhomogeneous INC-2 and homogeneous HNC-2 cases. Specifically, while the INC-2 and A-2 fields are in excellent agreement, the HNC-2 fields showed no signs of even approaching a steady state. Long after the INC-2 simulation had reached a steady state, the HNC-2 fields continued to amplify and develop asymmetric structures associated with flow nonlinearities. The very close agreement between the A-2 solution and the steady state in the INC-2 simulation is shown for the u field in Fig. 5. The u field in the disastrous HNC-2 simulation, at a time when a steady state had already been attained in the INC-2 simulation, is shown in Fig. 6.

441

442

443

444

425

426

427

428

429

430

431

432

433

434

435

436

437

438

439

440

4 Summary

The linearized Boussinesq equations for the motion of a viscous stably stratified fluid are solved analytically for a surface buoyancy that varies laterally as a square wave.

The solution describes two-dimensional laminar convective structures such as thermal convective rolls and updraft/downdraft pairs. The main applications of the solution may be in code verification and the evaluation of different implementations of the surface pressure condition for the pressure Poisson equation. Tests have been conducted for cases where the aspect ratios of the thermal disturbance have been large and small. With attention restricted to disturbances of sufficiently small amplitude, the linear solution and numerically simulated fields with the inhomogeneous Neumann condition for pressure (which is appropriate in the context of the particular fractional step procedure adopted in our DNS code) have been found to be in excellent agreement for both tests. However, in tests with a mis-specified Neumann condition, an excellent agreement with the analytical solution has been found only for the deep (small aspect ratio) disturbance case; errors in the shallow (large aspect ratio) disturbance case have been catastrophic.

445

446

447

448

449

450

451

452

453

454

455

456

458 Appendix A: Comment on the pressure condition at a lower solid surface

459 Consider a three-dimensional Boussinesq system with equation of motion,

460
$$\frac{\partial \mathbf{u}}{\partial t} = -\nabla \Pi + \nu \nabla^2 \mathbf{u} + \mathbf{F}. \tag{A1}$$

- 461 Here $\mathbf{u} = (u, v, w)$ is the three-dimensional velocity vector, Π is a kinematic pressure
- 462 perturbation, ν is the kinematic viscosity coefficient, and ${\bf F}$ is the sum of nonlinear
- acceleration and buoyancy terms. Applying the incompressibility condition,

$$\nabla \cdot \mathbf{u} = 0 \,, \tag{A2}$$

- in the equation that results from taking the divergence of (A1) (e.g., Orszag et al., 1986)
- 466 yields the Poisson equation,

$$\nabla^2 \Pi = \nabla \cdot \mathbf{F} \,. \tag{A3a}$$

- Although (A1) and (A2) imply (A3a), the reverse statement is not generally true.
- 469 Indeed, eliminating Π from between (A3a) and the equation arising from taking the
- 470 divergence of (A1) yields the diffusion equation $\partial \delta/\partial t = \nu \nabla^2 \delta$ for the velocity
- 471 divergence $\delta \equiv \nabla \cdot \mathbf{u}$, whose solution is (A2) only if δ is zero initially and on all
- 472 boundaries (Orszag et al., 1986; Gresho and Sani, 1987, Vreman 2014).
- 473 The same steps leading to (A3a) also lead to an alternative Poisson equation,

474
$$\nabla^2 \Pi = \nabla \cdot \left(\nu \nabla^2 \mathbf{u} + \mathbf{F} \right). \tag{A3b}$$

- 475 Although $\nabla \cdot \nu \nabla^2 \mathbf{u}$ was omitted in (A3a) [this term is zero if (A2) is satisfied], without
- 476 further constraints on δ (described above), (A2) may not be satisfied. Gresho and Sani

(1987) showed that the retention of $\nabla \cdot \nu \nabla^2 \mathbf{u}$ in (A3b) assures that (A2) is satisfied, 477 478 and thus leads to the paradox: "If you include it, you don't need it; if you don't include 479 it, you need it." Vreman (2014) revisited this paradox, and showed that for a standard 480 staggered method, the discretized form of (A3b) is equivalent to that of (A3a) supplemented with the constraint that $\nabla \cdot \nabla^2 \mathbf{u} = 0 \ (\nabla^2 \delta = 0)$ on points adjacent to the 481 482 solid boundary [with the same inhomogeneous Neumann boundary condition for Π implied for (A3a) and (A3b)]. When supplemented with this $\nabla^2 \delta = 0$ near-wall 483 condition, the diffusion equation for δ led to $\delta=0$ for all time. We note that (A3b) is 484 485 the form adopted in our numerical code.

Evaluating the vertical component of (A1) on the surface, where the impermeability condition applies, yields the inhomogeneous Neumann condition,

488
$$\frac{\partial \Pi}{\partial z}\Big|_{0} = \nu \frac{\partial^{2} w}{\partial z^{2}}\Big|_{0} + F_{z}\Big|_{0},$$
 (A4)

489

490

491

492

493

494

495

where $w \equiv \mathbf{k} \cdot \mathbf{u}$, $F_z \equiv \mathbf{k} \cdot \mathbf{F}$, \mathbf{k} is the upward unit vector, and () is a surface value. It has been argued that (A4), by itself, is not a proper boundary condition because it does not provide new information (it is not independent of the governing equations) and does not enforce the incompressibility condition (A2) at the boundary (Strikwerda, 1984; Henshaw, 1994; Sani et al., 2006). However, as pointed out by Henshaw (1994), many studies that impose (A4) (or a variant of it) also apply (A2) on the boundary.

In our numerical model, (A1) is integrated using a fractional step procedure with

explicit treatment of the viscous term. First, a provisional velocity field $\tilde{\mathbf{u}}$ that does not satisfy (A2) is obtained by integrating a discretized form of (A1) in which the pressure gradient is omitted. The provisional velocity is equal to the velocity at the end of the previous time step plus the sum of the forcing terms (nonlinear acceleration, buoyancy, and viscous stress) multiplied by the time step Δt . With the forcing terms explicitly evaluated, $\tilde{\mathbf{u}}$ is readily computed throughout the flow domain, including on the surface, where, in surface-forced flows, the buoyancy will make a substantial contribution. In terms of $\tilde{\mathbf{u}}$ and its vertical component \tilde{w} , (A3b) and (A4) become,

$$\nabla^2 \Pi = \frac{\nabla \cdot \tilde{\mathbf{u}}}{\Delta t},\tag{A5}$$

$$\frac{\partial \Pi}{\partial z}\Big|_{0} - \frac{1}{\Delta t} \tilde{w}\Big|_{0} = 0.$$
(A6)

In the second step, a velocity field that does satisfy (A2) is obtained by solving (A5) for Π and then adding the pressure gradient force associated with Π (multiplied by Δt) to $\tilde{\mathbf{u}}$.

In some explicit fractional step procedures (including the DNS code used in our study), the problem of solving (A5) subject to (A6) with $\tilde{\mathbf{u}}|_{0}$ evaluated from model data is replaced by what appears to be an entirely different (but is actually equivalent) problem: solving (A5) subject to the homogeneous Neumann condition,

$$\frac{\partial \Pi}{\partial z}\bigg|_{0} = 0, \tag{A7}$$

in concert with $\tilde{\mathbf{u}}\Big|_0$ being set to 0, obviating the need to calculate $\tilde{\mathbf{u}}\Big|_0$ from model data. It can be shown that $\tilde{w}|_{0}$ and the discretized form of $\partial \Pi/\partial z|_{0}$ appear in the discretized form of (A5) valid half a grid point above the physical surface as $\partial \Pi/\partial z|_0 - \tilde{w}|_0/\Delta t$, that is, in the same combination as they appear in (A6). Thus, setting $\tilde{w}|_{0}$ and $\partial \Pi/\partial z|_{0}$ to 0, is equivalent to implementing (A6) with the model-computed values of $\tilde{w}|_{0}$: the discretized form of (A5) near the surface is the same in either case. Moreover, on the C grid, setting the tangential components $\tilde{u}|_{0}$ and $\tilde{v}|_{0}$ to 0 only affects the values of \tilde{u} and \tilde{v} half a grid point beneath the physical boundary. These values do not appear in the discretized form of (A5) at any z-level, and thus have no bearing on the solution. In essence, the errors associated with the conflation of the two physically unjustifiable specifications (homogeneous Neumann condition for pressure, and $\tilde{\mathbf{u}}\Big|_{0} = 0$) cancel out. The homogeneous Neumann condition for pressure can be the source of confusion

514

515

516

517

518

519

520

521

522

523

524

525

526

527

528

529

530

531

if the context in which the condition is applied is not made clear: it would be a correct condition if $\tilde{\mathbf{u}}|_0$ is set to zero (per the equivalence described above), but it would be an incorrect condition if the explicit model-computed values of $\tilde{\mathbf{u}}|_0$ are used. In the experiments with the mis-specified condition described in Sect. 3, the homogeneous condition is imposed in the latter context. Unfortunately, in many numerical model descriptions, the nature of the surface pressure condition is left vague, for example, by

not indicating whether a Neumann condition is homogeneous or inhomogeneous, or, if a homogeneous Neumann condition is indicated, not mentioning how $\tilde{\mathbf{u}}|_{0}$ is treated.

Finally, we note that in fractional step procedures that treat the viscous term implicitly (e.g., Kim and Moin, 1985; Gresho, 1990; Armfield and Street, 2002; Guermond et al., 2006, and many others), the homogeneous Neumann condition is often applied as a surface condition for a Poisson equation, but it is again different from our implementation described in Sect. 3. In the implicit treatments, the provisional velocity is obtained as the solution of a boundary value problem ($\tilde{\mathbf{u}}\Big|_0$ should be specified; often it is set to 0) in which the relevant Poisson equation resembles (A5) but applies to a scalar function (sometimes called a pseudo-pressure) that is not the real pressure. Temam (1991) refers to this scalar as, "... a technical quantity, a mathematical auxiliary..." and advocates that it should not even be considered as an approximation of the pressure. Interestingly, in the context of implicit treatments, the homogeneous Neumann condition on the pseudo-pressure has sometimes been implicated as corrupting solution accuracy through the development of spurious numerical boundary layers adjacent to solid boundaries (Gresho, 1990; Guermond et al., 2006; Hosseini and Feng, 2011).

549

550

551

548

532

533

534

535

536

537

538

539

540

541

542

543

544

545

546

547

Code availability

The Fortran program used to generate output data files from the analytical solution is

available as a supplement to this article. That program (square.f) is configured for test A-1, but can be easily adjusted to run test A-2 or other tests. Running square.f automatically generates an output file for each dependent variable (e.g., u.dat) as well as an output file (square.out) that summarizes the test parameters and gives the computed values of the linearity ratios R_{η} and R_b defined in (3.1).

Acknowledgements. This research was supported by the National Science Foundation under Grant AGS-1359698. Comments by Chiel van Heerwaarden, Juan Pedro Mellado, Inanc Senocak, and an anonymous reviewer are gratefully acknowledged.

561 References

- 562 Armfield, S. and Street, R.: An analysis and comparison of the time accuracy of
- fractional-step methods for the Navier-Stokes equations on staggered grids. Int.
- J. Numer. Methods Fluids, 38, 255–282, 2002.
- Atkinson, B.: Meso-scale Atmospheric Circulations. Academic Press. 495 pp., 1981.
- Axelsen, S. L., Shapiro, A., and Fedorovich, E.: Analytical solution for katabatic flow
- induced by an isolated cold strip. Environ. Fluid Mech., 10, 387–414, 2010.
- 568 Briggs, G. A.: Surface inhomogeneity effects on convective diffusion. Boundary-Layer
- 569 Meteorol., 45, 117–135, 1988.
- 570 Chandrasekhar, S.: Hydrodynamic and Hydromagnetic Stability. Oxford University
- 571 Press. 652 pp., 1961.
- 572 Chorin, A. J.: Numerical solution of the Navier-Stokes equations. Math. Comput., 22,
- 573 745–762, 1968.
- Egger, J.: On the linear two-dimensional theory of thermally induced slope winds. Beitr.
- 575 Phys. Atmosph., 54, 465–481, 1981.
- 576 Fedorovich, E., Nieuwstadt, F. T. M., and Kaiser, R.: Numerical and laboratory study
- of a horizontally evolving convective boundary layer. Part I: Transition
- regimes and development of the mixed layer. J. Atmos. Sci., 58, 70–86, 2001.
- 579 Fedorovich, E. and Shapiro, A.: Structure of numerically simulated katabatic and
- anabatic flows along steep slopes. Acta Geophys., 57, 981–1010, 2009a.

- 581 Fedorovich, E. and Shapiro, A.: Turbulent natural convection along a vertical plate
- immersed in a stably stratified fluid. J. Fluid Mech., 636, 41–57, 2009b.
- 583 Gresho, P. M. and Sani, R. L.: On pressure boundary conditions for the incompressible
- Navier-Stokes equations. Int. J. Numer. Methods Fluids, 7, 1111–1145, 1987.
- 585 Gresho, P. M.: On the theory of semi-implicit projection methods for viscous
- incompressible flow and its implementation via a finite element method that also
- introduces a nearly consistent mass matrix. Part 1: Theory. Int. J. Numer.
- 588 Methods Fluids, 11, 587–620, 1990.
- 589 Guermond, J. L., Minev, P., and Shen, J.: An overview of projection methods for
- incompressible flows. Comput. Methods Appl. Mech. Engrg. 195, 6011–6045,
- 591 2006.
- Hadfield, M. G., Cotton, W. R., and Pielke, R. A.: Large-eddy simulations of thermally
- forced circulations in the convective boundary layer. Part I: A small-scale
- circulation with zero wind. Boundary-Layer Meteorol., 57, 79–114, 1991.
- Henshaw, W. D: A fourth-order accurate method for the incompressible Navier-Stokes
- equations on overlapping grids. J. Comput. Phys., 113, 13–25, 1994.
- 597 Hosseini, S. M. and Feng, J. J.: Pressure boundary conditions for computing
- incompressible flows with SPH. J. Comput. Phys., 230, 7473–7487, 2011.
- 599 Kang, S.-L., Lenschow, D., and Sullivan, P.: Effects of mesoscale surface thermal
- heterogeneity on low-level horizontal wind speeds. Boundary-Layer Meteorol.,

- 601 143, 409–432, 2012.
- 602 Kim, J. and Moin, P.: Application of a fractional-step method to incompressible
- 603 Navier-Stokes equations. J. Comput. Phys., 59, 308–323, 1985.
- 604 Kundu, P. K.: Fluid Mechanics. Academic Press. 638 pp., 1990.
- Mahrt, L., Sun, J., Vickers, D., MacPherson, J. I., Pederson, J. R., and Desjardins, R.
- L.: Observations of fluxes and inland breezes over a heterogeneous surface. J.
- 607 Atmos. Sci., 51, 2484–2499, 1994.
- 608 McPherson, R. A.: A review of vegetation-atmosphere interactions and their influences
- on mesoscale phenomena. Prog. Phys. Geog., 31, 261–285, 2007.
- Nordström, J., Mattsson, K., and Swanson, C.: Boundary conditions for a divergence
- free velocity-pressure formulation of the Navier-Stokes equations. J.
- 612 Comput. Phys., 225, 874–890, 2007.
- 613 Orszag, S. A., Israeli, M., and Deville, M. O.: Boundary conditions for incompressible
- 614 flows. J. Sci. Comput., 1, 75–111, 1986.
- 615 Patton, E. G., Sullivan, P. P., and Moeng, C.-H.: The influence of idealized
- heterogeneity on wet and dry planetary boundary layers coupled to the land
- 617 surface. J. Atmos. Sci., 62, 2078–2097, 2005.
- 618 Petersson, N. A.: Stability of pressure boundary conditions for Stokes and Navier-
- 619 Stokes equations. J. Comput. Phys., 172, 40–70, 2001.
- 620 Pielke, R. A.: Influence of the spatial distribution of vegetation and soils on the

- prediction of cumulus convective rainfall. Rev. Geophys., 39, 151–177, 2001.
- Rempfer, D.: On boundary conditions for incompressible Navier-Stokes problems. Appl.
- 623 Mech. Rev., 59, 107–125, 2006.
- 624 Segal, M. and Arritt, R. W.: Non-classical mesoscale circulations caused by surface
- sensible heat-flux gradients. Bull. Amer. Meteorol. Soc., 73, 1593–1604, 1992.
- 626 Shirokoff, D. and Rosales, R. R.: An efficient method for the incompressible Navier-
- Stokes equations on irregular domains with no-slip boundary conditions, high
- order up to the boundary. J. Comput. Phys., 230, 8619–8646, 2011.
- 629 Shapiro, A. and Fedorovich, E.: Similarity models for unsteady free convection flows
- along a differentially cooled horizontal surface. J. Fluid Mech., 736, 444–463,
- 631 2013.
- 632 Shapiro, A. and Fedorovich, E.: A boundary-layer scaling for turbulent katabatic flow.
- 633 Boundary-Layer Meteorol., 153, 1–17, 2014.
- 634 Simpson, J. E.: Sea Breeze and Local Winds. Cambridge University Press. 234 pp.,
- 635 1994.
- 636 Temam, R.: Remark on the pressure boundary condition for the projection method.
- 637 Theoret. Comput. Fluid Dynamics, 3, 181–184, 1991.
- 638 Strikwerda, J. C.: Finite difference methods for the Stokes and Navier-Stokes equations.
- 639 SIAM J. Sci. Stat. Comput., 5, 56–68, 1984.
- 640 van Heerwaarden, C. C., Mellado, J. P., and de Lozar, A.: Scaling laws for the

641	heterogeneously heated free convective boundary layer. J. Atmos. Sci., 71, 3975–
642	4000, 2014.
643	Vreman, A. W.: The projection method for the incompressible Navier-Stokes equations:
644	The pressure near a no-slip wall. J. Comput. Phys., 263, 353–374, 2014.
645	

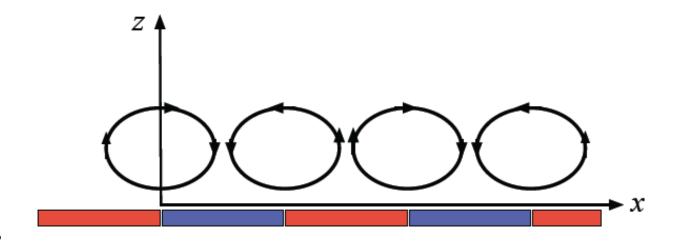


Figure 1. Schematic of two-dimensional (x, z) thermal convection induced by a surface buoyancy that varies horizontally (x) as a square wave. Red denotes positive surface buoyancy, blue denotes negative surface buoyancy.

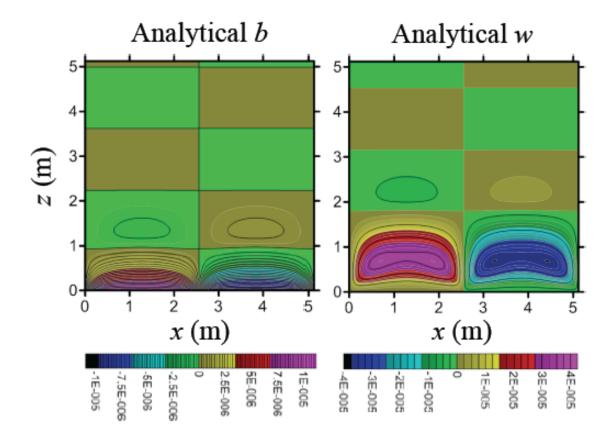


Figure 2. Vertical cross section of the analytical (A-1) buoyancy b and vertical velocity w fields from the first test case. Color bar units are m s⁻² for b, and m s⁻¹ for w.

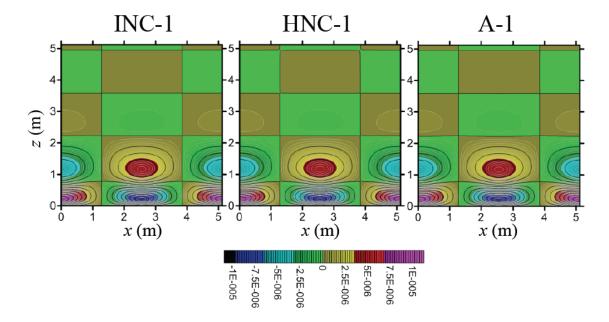


Figure 3. Vertical cross section of u from the first test case. A-1 is the analytical solution. INC-1 is the numerical simulation with inhomogeneous Neumann condition pressure. HNC-1 is the numerical simulation with the homogeneous Neumann condition for pressure. Color bar units are m s⁻¹.



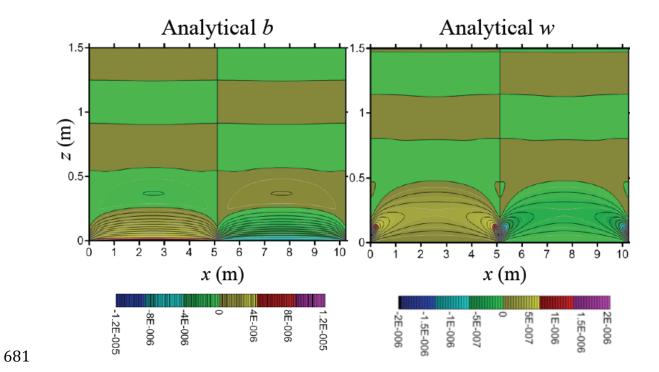


Figure 4. Vertical cross section of the analytical (A-2) buoyancy b and vertical velocity w fields from the second test case. Color bar units are m s⁻² for b, and m s⁻¹ for w.



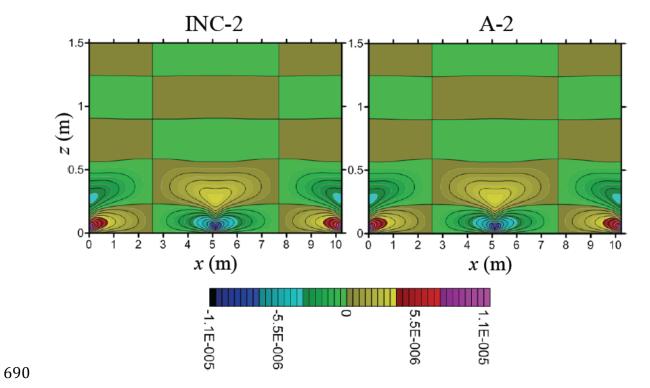


Figure 5. Vertical cross section of u from the second test case. A-2 is the analytical solution. INC-2 is the numerical simulation with inhomogeneous Neumann condition for pressure. Color bar units are m s⁻¹.

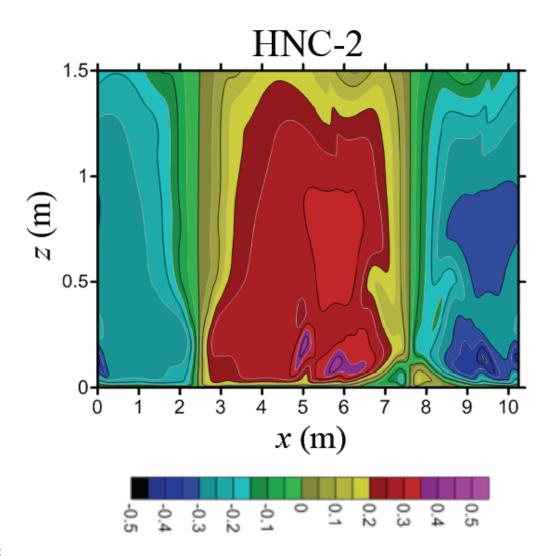


Figure 6. Vertical cross section of u from HNC-2, the numerical simulation with 701 homogeneous Neumann condition for pressure in the second test case. Color bar units 702 are m s⁻¹.