1	Sensitivity of Chemistry-Transport Model Simulations to the Duration of Chemical and
2	Transport Operators: A Case Study with GEOS-Chem v10-01
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17	Abstract
18	Chemistry-transport models involve considerable computational expense. Fine temporal
19	resolution offers accuracy at the expense of computation time. Assessment is needed of the
20	sensitivity of simulation accuracy to the duration of chemical and transport operators. We conduct
21	a series of simulations with the GEOS-Chem chemistry-transport model at different temporal and
22	spatial resolutions to examine the sensitivity of simulated atmospheric composition to operator
23	duration. Subsequently, we compare the species simulated with operator durations from 10 min to

24 60 min as typically used by global chemistry-transport models, and identify the operator durations 25 that optimize both computational expense and simulation accuracy. We find that longer continuous 26 transport operator duration increases concentrations of emitted species such as nitrogen oxides and 27 carbon monoxide since a more homogeneous distribution reduces loss through chemical reactions 28 and dry deposition. The increased concentrations of ozone precursors increase ozone production 29 with longer transport operator duration. Longer chemical operator duration decreases sulfate and 30 ammonium but increases nitrate due to feedbacks with in-cloud sulfur dioxide oxidation and 31 aerosol thermodynamics. The simulation duration decreases by up to a factor of 5 from fine (5 32 min) to coarse (60 min) operator duration. We assess the change in simulation accuracy with 33 resolution by comparing the root mean square difference in ground-level concentrations of 34 nitrogen oxides, secondary inorganic aerosols, ozone and carbon monoxide with a finer temporal 35 or spatial resolution taken as "truth". Relative simulation error for these species increases by more 36 than a factor of 5 from the shortest (5 min) to longest (60 min) operator duration. Chemical operator 37 duration twice that of the transport operator duration offers more simulation accuracy per unit 38 computation. However, relative simulation error from coarser spatial resolution generally exceeds that from longer operator duration; e.g. degrading from 2° x 2.5° to 4° x 5° increases error by an 39 40 order of magnitude. We recommend prioritizing fine spatial resolution before considering different 41 operator durations in offline chemistry-transport models. We encourage chemistry-transport model 42 users to specify in publications the durations of operators due to their effects on simulation accuracy. 43

## 44 **1 Introduction**

Global and regional chemistry-transport models (CTMs) have a wide range of applications in
studies of climate, air quality, and biogeochemical cycling. The last few decades have witnessed

47 rapid development of modeling sophistication to tackle these issues, but that development is 48 associated with increasing computational expense. Typically, Eulerian models divide the 49 atmosphere into numerous (10<sup>4</sup>-10<sup>8</sup>) grid boxes and solve the mass continuity equation to simulate 50 atmospheric composition. Numerical solution of the mass continuity equation involves separating 51 the different chemical and transport processes (or operators) through operator splitting. The 52 concentrations of simulated species are sensitive to the duration of operators used in the CTM. 53 Attention is needed to understand how operator duration affects model performance.

54 Numerous studies have examined the sensitivity of simulations to grid resolution for ozone (Jang 55 et al., 1995; Esler et al., 2004; Ito et al., 2009; Yu et al., 2016), ozone production efficiency (Liang 56 and Jacobson 2000), and ozone sensitivity to precursor emissions (Cohan et al., 2006; Henderson 57 et al., 2010). Simulation error increases proportional to the size of the horizontal grid (Wild and 58 Prather, 2006; Prather et al., 2008). Biases can be reduced by simulating sub grid scale processes 59 such as emission plumes from point sources (Sillman et al., 1990; Valin et al., 2011), aircraft 60 exhaust (Kraabøl et al., 2002), ship exhaust (Vinken et al., 2011), mineral dust emissions (Ridley 61 et al., 2013), and lightning (Cooper et al., 2014). The spatial and temporal resolution of the 62 meteorological fields used in CTMs can also influence model processes (Bian et al., 2009). The 63 spatiotemporal variation of carbon monoxide is better represented with finer grid resolution (Wang 64 et al., 2004; Chen et al., 2009; Yan et al., 2014). Moreover, fine horizontal resolution is important 65 for air quality exposure assessment and health impact studies (Punger and West, 2013; Fountoukis 66 et al., 2013; Thompson et al., 2014; Li et al., 2015). Fine vertical resolution can better represent 67 the effects of convection (Rind et al., 2007; Arteta et al., 2009). Simulations are also sensitive to 68 operator durations (Mallet et al., 2007; Santillana et al., 2016), however, few studies have examined this sensitivity. 69

70 CTMs solve the continuity equation for tens to hundreds of chemical species, each with number
71 density *n*, for individual grid boxes defined in the Eulerian model.

72

$$\frac{\partial n}{\partial t} = -\nabla \cdot nU + P - L \tag{1}$$

73  $\partial n/\partial t$  represents the local temporal evolution of  $n_{.} - \nabla \cdot nU$  represents the transport flux divergence term, where U is the wind velocity vector. P and L are the local production and loss terms 74 75 respectively. Typically, the above equation is discretized in space, and the continuity equation is 76 simulated as a system of coupled non-linear partial differential equations with chemical and 77 transport operators. These chemical and transport operators are usually simulated sequentially 78 through operator splitting to increase computational efficiency (Hundsdorfer and Verwer, 2003). 79 The transport operator involves solving the 3-D advection equation using efficient numerical 80 schemes (Prather, 1986; Lin and Rood, 1996). Boundary layer mixing, convection, emission and 81 deposition are often simulated as individual operators. The chemical operator representing the 82 temporal evolution of local sources and sinks involves numerically solving a system of coupled 83 ordinary differential equations using efficient solvers (Jacobson and Turco, 1994; Damian et al., 84 2002). The integration timestep in a differential equation solver is important for efficient and 85 accurate solution (Jacobson and Turco, 1994). Moreover, the model accuracy is affected by the 86 duration of chemical and transport operators (Mallet and Sportisse, 2006; Mallet et al., 2007), and 87 the order in which these operators are applied (Sportisse, 2000; Santillana et al., 2016). The 88 operator splitting method requires the coupling between individual operators to be negligible over 89 the operator duration. However, reducing operator durations increases computational expense. 90 Attention is needed to this tradeoff.

We examine the sensitivity of a CTM to operator duration by conducting a series of simulations at
 different horizontal resolutions and operator durations. We then identify the optimal operator

duration from the range of operator durations from 10 min to 60 min usually used by global CTMs
(e.g., Horowitz et al., 2003; Huijnen et al. 2010). Section 2 describes the sensitivity simulations,
the method to quantify the simulation error, as well as the method to identify the simulation
operator durations that best account for both computational expense and simulation accuracy.
Comparison of the sensitivity simulations, description of resolution-dependent errors, and the
identification of appropriate chemical and transport operator durations are examined in section 3.

#### 99 2 Materials and Methods

## 100 2.1 GEOS-Chem simulations

101 We conduct a series of sensitivity simulations with the GEOS-Chem CTM (version 10-01; 102 www.geos-chem.org) at different horizontal resolutions and operator durations to examine the 103 individual sensitivities to chemical and transport operator durations. The GEOS-Chem model (Bey 104 et al., 2001) is used by about 100 research groups worldwide to simulate the oxidant-aerosol 105 system. GEOS-Chem has the capability to be driven with several generations of assimilated 106 meteorological data from the Goddard Earth Observing System (GEOS) at the NASA Global 107 Modeling Assimilation Office (GMAO). For computational expedience, GEOS-Chem global 108 simulations are often conducted using horizontal resolutions of either 4° x 5° or 2° x 2.5° degraded 109 from the native resolution of GEOS meteorology. GEOS-Chem also has the capability for nested 110 regional simulations where the global model provides dynamic boundary condition to the finer 111 regional grids (Wang et al., 2004; Chen et al., 2009; Zhang et al., 2011; van Donkelaar et al., 2012). We use the GEOS-5.2.0 meteorology available at a native horizontal resolution of 0.5° x 0.667° 112 113 (Rienecker et al., 2008). It includes three-hour averaged 2-D fields such as mixed layer depth, and 114 six-hour averaged 3-D fields such as zonal and meridional wind, and convective mass flux. The

height of the lowest level of the model is approximately 130 meters above the sea level, with 47vertical levels.

GEOS-Chem performs species advection (A), vertical mixing (V), cloud convection (Z) and wet
deposition (W) for every transport operator duration (T), as well as dry deposition (D), emissions
(E), and chemistry (G) for every chemical operator duration (C) in the following order,

$$A(T) \cdot D(C) \cdot E(C) \cdot V(T) \cdot Z(T) \cdot G(C) \cdot W(T)$$
(2)

121 The traditional transport operator durations are 30 minutes at 4° x 5° resolution, 15 minutes at 2° x 2.5° resolution, and 10 min at 0.5° x 0.667° resolution. The traditional chemical operator duration 122 123 is set to either 60 min or twice the transport operator duration based on the Strang operator splitting 124 scheme (Strang, 1968) which follows  $T \cdot C \cdot T \cdot T \cdot C \cdot T$  order repetitively with  $C = 2 \times T$ . 125 Transport operations are repeated twice before a chemical operation when  $C = 2 \times T$ . We also 126 consider an alternate splitting scheme which follows  $T \cdot C \cdot T \cdot C$  order repetitively with C = T. Changes in operator duration from  $C = 2 \times T$  to C = T include effects of both time truncation (T  $\cdot$ 127 128 T to T) and operator splitting.

129 Advection is based on the multi-dimensional flux-form semi-Lagrangian advection scheme (Lin 130 and Rood, 1996; Lin et al., 1994), with an additional pressure-fixer algorithm implemented for the conservation of species mass (Rotman et al., 2004). The cloud convection operator couples 131 132 transport by convection (Balkanski et al., 1993; Wu et al., 2007) with gas-aerosol wet deposition 133 (Liu et al., 2001; Wang et al., 2011; Amos et al., 2012). GEOS-Chem uses an internal integration 134 timestep of 5 min for convective mixing within the cloud convection operator. The wet deposition 135 operator includes scavenging by large-scale precipitation through first order operators, rainout and 136 washout (Balkanski et al., 1993). We use a non-local boundary layer mixing scheme for vertical

transport (Holtslag and Boville, 1993, Lin and McElroy, 2010). Emissions are processed through
the HEMCO module (Keller et al., 2014). A resistance-in-series method is used for dry deposition
of species (Wesely, 1989; Wang et al., 1998; Zhang et al., 2001; Fisher et al., 2011).

140 GEOS-Chem uses a Sparse Matrix Vectorized GEAR II chemistry solver (Jacobson and Turco, 141 1994; Jacobson, 1995; 1998). The oxidant-aerosol chemistry simulation includes organic and black 142 carbon (Park et al., 2003), mineral dust (Fairlie et al., 2007; Zender et al., 2003; Ginoux et al., 143 2001), sea salt (Alexander et al., 2005; Jaegle et al., 2011), and the sulfate-nitrate-ammonium 144 system (Park et al., 2004). The photolysis frequency is calculated (Mao et al., 2010; Eastham et 145 al., 2014) at the middle of the chemical operator duration using the Fast-JX algorithm (Bian and 146 Prather, 2002). Simulation of gas-aerosol interactions are performed within the chemistry operator 147 by aerosol extinction effects on photolysis rates (Martin et al., 2003), and heterogeneous chemistry 148 (Jacob, 2000) including aerosol uptake of N<sub>2</sub>O<sub>5</sub> (Evans and Jacob, 2005) and HO<sub>2</sub> (Mao et al., 149 2013). The ISORROPIA II thermodynamic module (Fontoukis and Nenes, 2007) performs 150 aerosol-gas partitioning (Pye et al., 2009).

We conduct simulations for 2010 July at two horizontal resolutions of  $4^{\circ} \times 5^{\circ}$  and  $2^{\circ} \times 2.5^{\circ}$  globally, and  $0.5^{\circ} \times 0.667^{\circ}$  over the North America ( $140^{\circ}W-40^{\circ}W$ ,  $10^{\circ}N-70^{\circ}N$ ) and East Asia ( $70^{\circ}W-153$   $150^{\circ}W$ ,  $11^{\circ}S-55^{\circ}N$ ) nested regions. We use the  $4^{\circ} \times 5^{\circ}$  global simulation to archive dynamic boundary conditions every three hours for the nested simulations. We use one month spin up with each GEOS-Chem simulation to reduce the influence of initial conditions.

## 156 **2.2 Computing platform**

157 We conduct all simulations on the same computing platform to compare their computational 158 performance. We use the Glooscap cluster of the Atlantic Computational Excellence Network (ACENET) Consortium of Canadian Universities (http://www.ace-net.ca/wiki/Glooscap). The
operating system is Linux 4.8. We use Intel Fortran compiler version 12. Each GEOS-Chem
simulation is submitted as a 16-thread parallelized job on a single node.

We calculate the CPU time for the month of July for each operator separately using the Fortranintrinsic routine, CPU\_TIME. We found this value identical to the one calculated using the Linux command 'qacct –j'. To reduce the effects of other jobs on the shared cluster, we repeat simulations five times, while excluding data output operations to minimize sensitivity to system input/output, and use the median to represent CPU time. We also report the standard error over the five simulations.

## 168 **2.3 Assessing the relative simulation error**

169 We treat the simulation with the shortest operator duration as the most accurate. This approach 170 exploits the reduction in error associated with coupling across operators as operator duration 171 diminishes. Assessing simulation error versus operator duration through comparison with 172 observations is impaired by imperfect model processes, by the sparseness of measurements, and by model-observation representativeness biases. We take as "truth" the concentrations simulated 173 174 with a chemical operator duration (C) of 10 minutes and a transport operator duration (T) of 5 175 minutes (represented as C10T05). Finer resolutions are computationally prohibitive. We define the relative simulation error  $E_{sim}^{s}$  for species s as the root mean square error (RMSE) of the species 176 177 concentrations simulated with the finest resolution ("truth") and the simulation under consideration 178 (Sim), normalized by the concentrations in simulation "truth",

179  
$$E_{sim}^{s} = \frac{\sqrt{N}\sqrt{\sum_{i=1}^{i=N} (Truth_{i}^{s} - Sim_{i}^{s})^{2}}}{\sum_{i=1}^{i=N} (Truth_{i}^{s} - Sim_{i}^{s})^{2}}$$

$$\sum_{i=1}^{i=N} Truth_i^s$$
(3)

181 where, *i* represents a particular grid box, with a total number of *N* grid boxes of interest. RMSE in 182 the numerator is chosen instead of absolute difference to more heavily penalize extrema. 183 Normalization with the mass of the "true" simulation is intended to cross-compare  $E_{sim}^{s}$  of different 184 species.  $E_{sim}^{s}$  captures the variation of a species *s* from the "true" simulation.

185 We focus on four key species relevant to atmospheric chemistry, namely nitrogen oxides (NO<sub>x</sub> = 186  $NO + NO_2$ ), secondary inorganic aerosols (SIA: sum of sulfate, nitrate and ammonium), ozone 187  $(O_3)$ , and carbon monoxide (CO). These species represent a range of lifetimes from a day  $(NO_x)$ 188 to weeks (CO). The focus on SIA is designed to devote more attention to chemically active species 189 than to mineral dust and sea salt. We sample the instantaneous values of simulated ground-level 190 concentrations of these atmospheric species every 60 min to span the diurnal variation of chemical 191 environments. We focus on concentrations in July near the Earth's surface when and where 192 chemical and transport timescales tend to be short.

#### 193

#### 2.4 Identifying the optimal operator duration

A practical way to select optimal chemical and transport operator durations is to identify the simulation with the lowest error ( $E_{sim}^s$ ) per unit of computation time. To quantify the simulation accuracy per unit CPU time, we propose a simple metric, the CPU-time adjusted Composite Normalized Error (CNE) which represents a tradeoff between the simulation accuracy, and the associated computation expense. This is performed by normalizing the relative simulation error  $E_{sim}^s$  for species *s* by the CPU time *t* for the simulation under consideration  $t_{sim}$  and for a reference simulation  $t_{ref}$ , and taking the mean of the four species.

202 
$$CNE = \left(\frac{1}{4} \times \sum_{s} \frac{E_{sim}^{s}}{E_{ref}^{s}}\right) \times \left(\frac{t_{sim}}{t_{ref}}\right)$$
(4)

We normalize  $E_{sim}^{s}$  by the reference  $E_{ref}^{s}$  so that the CPU-time adjusted Composite Normalized Error for each species is of similar magnitude. The variation of CNE across operator durations is unaffected by the choice of reference simulation; C10T10 used here. The relative value of CPU time versus simulation accuracy is subjective and depends on scientific objective. This definition of CNE gives equal weighting to the respective cost of CPU time and simulation accuracy. The simulation with the lowest CNE is used to identify an optimal chemical and transport operator duration.

#### 210 **3 Results and discussion**

211 Figure 1 shows the computational performance for the series of GEOS-Chem simulations 212 conducted here. The CPU time decreases by factors of 3-5 from fine to coarse operator duration. 213 The CPU time increases by about a factor of 4 from 4° x 5° to 2° x 2.5° and another factor of 2 to a single nested simulation at 0.5° x 0.667°. The linearity from 4° x 5° to 2° x 2.5° implies that grid 214 215 boxes are sufficiently large that CPU time is proportional to the number of grid boxes, and that 216 transport integration timesteps constrained by the Courant-Freidrich-Lewy criterion (Courant et 217 al., 1967) are largely unaffected by changes to grid box size at these resolutions. Comparison of 218 individual CPU times for chemical and transport operators shows that performing a single cycle 219 of all chemical operations takes ~4 times that of a single cycle of transport operations at the global 220 scale. This factor is reduced for nested simulations due in part to the additional CPU time for 221 simulating boundary conditions.

Figure 2 illustrates the sensitivity of the simulations to chemical and transport operators at 2° x 2.5° horizontal resolution. The left column shows the species concentrations for the "true"

224 simulation (C10T05). The middle column shows the difference in species concentrations from 225 doubling the transport operator duration. This doubling is in practice a change in time truncation 226 of the transport operator from  $T \cdot C \cdot T \cdot T \cdot C \cdot T$  to  $T \cdot C \cdot T \cdot C$  since the transport operator must 227 keep pace with the chemistry operator. Increasing the transport operator duration tends to increase 228 concentrations of emitted species like CO and NO<sub>x</sub> over source regions since species are more 229 uniformly mixed by long continuous operator durations before loss processes such as dry 230 deposition and chemistry occur. More homogeneous fields have lower dry deposition rates as a 231 larger fraction is mixed aloft, and lower chemical loss rates depending on the chemical regime. 232 The increase in CO over source regions is partly associated with decreases in OH. Increasing 233 concentrations of ozone precursors increases ozone production ( $P[O_3]$ ). Wild and Prather (2006) 234 similarly found that ozone production increases at coarser horizontal resolution. Increasing the 235 transport operator duration increases SIA components, especially over the source regions of East 236 Asia, North India, and North America.

237 The right column in Fig. 2 shows the change in species concentrations from increasing the 238 chemical operator duration. Hydroxyl radical concentrations increase, NO<sub>x</sub> concentrations 239 decrease, and P[O<sub>3</sub>] decreases with increasing chemical operator durations over source regions. 240 Berntsen and Isaken (1997) found that the error introduced by coarser chemical operator durations 241 is higher in polluted regions than the clean background due to the increased time lag, and invariant 242 production and loss across rapid chemical cycles. A longer chemical operator duration decreases 243 sulfate and ammonium but increases nitrate over source regions. Inspection of SO<sub>2</sub> and H<sub>2</sub>O<sub>2</sub> fields 244 indicates that sulfate formation through H<sub>2</sub>O<sub>2</sub> in clouds decreases at longer chemical operator 245 durations. In turn, SO<sub>2</sub> and NH<sub>3</sub> concentrations increase at longer chemical operator durations due 246 to the corresponding decreases in ammonium sulfate or ammonium bisulfate. The additional free

ammonia at longer chemical operator durations tends to promote regional ammonium nitrate formation depending on local thermodynamics. An increase of total SIA mass with increasing chemical operator duration is driven by nitrate and ammonium, and partially compensated by a reduction in sulfate, especially downwind of source regions. We find similar spatial patterns for other operator duration combinations, and other horizontal resolutions.

Figure 3 shows the sensitivity of simulated species to changes in operator duration (C20T10 to C10T05) at two other horizontal resolutions (global 4° x 5°, and nested North America 0.5° x 0.67° simulations) considered here. Spatial patterns of monthly mean ground-level concentrations, and absolute differences are similar, albeit with finer spatial heterogeneity resolved in the nested simulation. However, some resolution dependent differences do arise reflecting nonlinear feedbacks.

258 Figures 4 shows the relative simulation error for nitrogen oxides, secondary inorganic aerosols, 259 ozone and carbon monoxide with varying operator durations at 2° x 2.5° horizontal resolution. 260 Relative simulation errors for all these major species increase by more than a factor of 5 from the 261 shortest to longest operator duration. Errors increase fairly smoothly with increasing chemical and 262 transport operator duration until the transport operator duration exceeds 30 min. Then errors 263 increase by an order of magnitude for long lived species of  $O_3$  and CO. The saw-tooth pattern for 264 CO versus O<sub>3</sub> reflects a greater sensitivity of CO to transport operator duration and a greater 265 sensitivity of O<sub>3</sub> to chemical operator duration. Relative simulation errors for other horizontal 266 resolutions follow similar pattern. These relative errors of 5% - 35% for NOx and SIA are 267 comparable to typical model-observation errors of  $\sim 30\%$  for NO<sub>x</sub> (Boersma et al., 2008; Hudman 268 et al., 2006) and 20 - 40% for SIA (Philip et al., 2014; Heald et al., 2012). Operator duration errors

of <2% for O<sub>3</sub> and CO are smaller than typical model-observation errors of  $\sim20\%$  for ozone (Zhang et al., 2011; Wang et al., 2009) and 10 - 20% for CO (Duncan et al., 2007; Shindell et al., 2006).

271 Figure 5 shows the difference in simulated species at 2° x 2.5° horizontal resolution for the GEOS-272 Chem traditional (C30T15) minus the finest operator durations considered (C10T05). The spatial 273 variation for the monthly mean ground-level concentrations is generally within 5-15% for short 274 lived species like NO<sub>x</sub> and SIA, and within 1% for longer lived species like O<sub>3</sub> and CO. Santillana 275 et al. (2016) similarly found an upper limit of 10% for operator splitting errors. However, the 276 maximum hourly spatial variation can exceed 50% for short lived species and 5% for longer lived 277 species. The spatial pattern of extrema resembles that of the monthly mean, albeit with more 278 heterogeneity from synoptic variation.

We also examined the diurnal variation and vertical profile of extrema. Extrema arise from all times of day with a slight tendency for larger values for  $NO_x$  at night, for  $O_3$  near sunrise and sunset, and for SIA and CO near noon. Zonal mean vertical profiles exhibit largest differences in the lower troposphere for  $NO_x$  and SIA, with more homogeneous differences throughout the troposphere for  $O_3$  and CO. Near the subtropical jets of the upper troposphere  $O_3$  and CO have maximum extrema of up to 3%.

Figure 6 shows the CPU-time adjusted Composite Normalized Error for the GEOS-Chem simulations at various horizontal resolutions and operator durations. The CNE is significantly higher with C = T than  $C = 2 \times T$ . We confirmed this tendency with different choices of "truth" (such as C05T05, C10T10) or reference (such as C10T05) simulations. This finding motivates the traditional approach of using  $C = 2 \times T$  in GEOS-Chem simulations. Applying the chemical operator as frequently as the transport operator (with C = T) appears to increase computation cost

291 with little benefit in accuracy. The CNE for all three horizontal resolutions have noisy minima 292 with a chemical operator duration of 20 min and a transport operator duration of 10 min (C20T10). 293 A unit of computation time has a similar efficiency for a small range of operator durations from 294 10 min to 20 min. We found similar patterns in the variation of CNE with operator durations with 295 CNE calculated for selected domains, such as over Northern Hemisphere, nested model regions, land grid boxes, and over the entire troposphere. We conducted additional simulations at 4° x 5° 296 297 horizontal resolution for January 2011 with a spin up of 7 months, and found similar patterns in 298 CNE.

The relative simulation error decreases by 40-50% (Fig. 4) by changing the operator duration from the traditional (C30T15) to the optimal (C20T10) at  $2^{\circ} \times 2.5^{\circ}$  horizontal resolution. The relative spatial variations are <20% for NO<sub>x</sub> and SIA, and <1% for O<sub>3</sub> and CO. However, the CPU time increases by 20% by the decrease in operator duration.

Table 1 shows the relative simulation error at 4° x 5° horizontal resolution with "truth" at 2° x 2.5° 303 304 resolution (C10T05) to investigate the tradeoff between horizontal resolution and operator 305 duration. The simulation error for all species at 4° x 5° resolution increases by an order of 306 magnitude compared to 2° x 2.5° resolution for any choice of operator duration tested here. The 307 error in this configuration is insensitive to operator duration, and dominated by representativeness differences due to spatial structure resolved at 2° x 2.5° resolution, but not at 4° x 5° resolution. 308 309 Nonlinear chemistry at different horizontal resolutions (e.g., Wild and Prather, 2006) also plays a 310 role. Numerical errors due to advection processes generally exceed those from operator splitting 311 (e.g., Prather et al., 2008; Santillana et al., 2016). We therefore recommend prioritizing horizontal 312 resolution over operator duration for offline CTMs using time-averaged meteorological fields as 313 tested here. As meteorological fields used in CTMs become available at finer temporal and spatial 314 resolution, the value of shorter operator duration should further increase. We encourage CTM users

315 to specify in publication the duration of operators due to its effect on simulation accuracy.

## 316 4 Conclusions

317 The computational expense of chemistry-transport models warrants investigation into their 318 efficiency and accuracy. Solving the continuity equation in CTMs through operator splitting 319 method offers numerical efficiency, however, few studies have examined the implications of 320 operator duration on simulation accuracy. We conducted simulations with the GEOS-Chem model 321 for multiple choices of operator duration from 10 min to 60 min as typically used by global CTMs. 322 We found that longer continuous transport operator durations increase ozone precursors and ozone 323 production over source regions since a more homogeneous distribution reduces loss through 324 chemical reactions and dry deposition. Longer chemical operator durations decrease NO<sub>x</sub> and 325 ozone production over source regions. Longer chemical operator durations reduce sulfate and 326 ammonium concentrations, however increase nitrate due to feedbacks with in-cloud SO<sub>2</sub> oxidation 327 and local aerosol thermodynamics.

328 We investigated the computational efficiency with the GEOS-Chem model, and found that the 329 simulation computation time decreases by up to a factor of 5 from short (C10T05) to long 330 (C60T60) operator duration. The chemical operator consumes about four times the CPU time of 331 the transport operator. We subsequently compared the root mean square differences in the ground-332 level concentrations of nitrogen oxides, secondary inorganic aerosols (SIA), ozone and carbon 333 monoxide with a finer temporal or spatial resolution taken as "truth", and estimated the relative 334 simulation error. The relative simulation error for these species increases by more than a factor of 335 5 from the shortest to longest operator duration. Monthly mean simulation errors of about 30% for

 $NO_x$  and SIA from long operator duration are comparable to typical model-observation errors,

337 while simulation errors for CO and  $O_3$  tend to be less than 2% for operator duration < 30 min.

338 In order to account for simulation accuracy with computational cost, we proposed a metric, CPU-339 time adjusted Composite Normalized Error that identifies the operator duration with respect to 340 CPU cost. We find greater efficiency of using  $C = 2 \times T$  than C = T for all horizontal resolutions. 341 The Composite Normalized Error exhibits a noisy minimum for a chemical operator duration of 342 20 min and transport operator duration of 10 min for the range of operator durations and horizontal 343 resolutions considered here. Nonetheless, the relative simulation error from changing horizontal 344 resolution exceeds that from changing operator durations within a horizontal resolution. We 345 recommend prioritizing fine spatial resolution before considering different operator durations in 346 offline CTMs with time-averaged archived meteorological fields as tested here. The importance 347 of shorter operator durations should increase with the availability of time-averaged meteorological 348 fields at higher temporal resolution. Short operator durations could offer even greater benefits to 349 simulation accuracy in online CTMs that offer meteorological fields at temporal resolutions closer 350 to operator duration. We encourage CTM users to specify in publications the durations of operators 351 due to their effects on simulation accuracy.

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## 357 Code Availability

- 358 The GEOS-Chem code is freely accessible to the public, by following the guidelines in
- 359 <u>http://wiki.geos-chem.org/</u>). This work used GEOS-Chem version 10-01.

360

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## 367 References

- Alexander, B., Park, R. J., Jacob, D. J., Li, Q. B., Yantosca, R. M., Savarino, J., Lee, C. C. W.
- and Thiemens, M. H.: Sulfate formation in sea-salt aerosols: Constraints from oxygen isotopes, J.
- Geophys. Res., 110, D10307, doi:10.1029/2004JD005659, 2005.
  371
- Amos, H. M., Jacob, D. J., Holmes, C. D., Fisher, J. A., Wang, Q., Yantosca, R. M., Corbitt, E.
- 373 S., Galarneau, E., Rutter, A. P., Gustin, M. S., Steffen, A., Schauer, J. J., Graydon, J. A., St
- 374 Louis, V. L., Talbot, R. W., Edgerton, E. S., Zhang, Y., and Sunderland, E. M.: Gas-particle
- partitioning of atmospheric Hg(II) and its effect on global mercury deposition, Atmos. Chem.
  Phys., 12, 591-603, doi:10.5194/acp-12-591-2012, 2012.
- 377
- Arteta, J., Marecal, V., and Riviere, E. D.: Regional modelling of tracer transport by tropical
  convection Part 2: Sensitivity to model resolutions, Atmos. Chem. Phys., 9, 7101-7114,
  doi:10.5194/acp-9-7101-2009, 2009.
- 381
- Balkanski, Y. J., Jacob, D. J., Gardner, G. M., Graustein, W. C., and Turekian, K. K.: Transport
   and residence times of tropospheric aerosols inferred from a global three-dimensional simulation
   of <sup>210</sup>Pb, J. Geophys. Res., 98, 20573-20586, doi:10.1029/93JD02456, 1993.
- 385
- 386
- 387

- Berntsen, T. K., and Isaksen, I. S. A.: A global three-dimensional chemical transport model for
- the troposphere: 1. Model description and CO and ozone results, J. Geophys. Res., 102, 2123921280, doi:10.1029/97JD01140, 1997.
- 391
- Bey, I., Jacob, D. J., Yantosca, R. M., Logan, J. A., Field, B. D., Fiore, A. M., Li, Q. B., Liu, H.
- 393 G. Y., Mickley, L. J., and Schultz, M. G.: Global modeling of tropospheric chemistry with 394 assimilated meteorology: Model description and evaluation, J. Geophys. Res., 106, 23073-
- 395 23095, doi:10.1029/2001JD000807, 2001.
- 396
- Bian, H., Chin, M., Rodriguez, J. M., Yu, H., Penner, J. E., and Strahan, S.: Sensitivity of aerosol
  optical thickness and aerosol direct radiative effect to relative humidity, Atmos. Chem. Phys., 9,
  2375-2386, doi:10.5194/acp-9-2375-2009, 2009.
- 400
- Bian, H., and Prather, M. J.: Fast-J2: Accurate simulation of stratospheric photolysis in global
- 402 chemical models, J. Atmos. Chem., 41, 281-296, doi:10.1023/A:1014980619462, 2002. 403
- 404 Boersma, K. F., Jacob, D. J., Eskes, H. J., Pinder, R. W., Wang, J., and van der A, R. J.:
- 405 Intercomparison of SCIAMACHY and OMI tropospheric NO<sub>2</sub> columns: Observing the diurnal
- 406 evolution of chemistry and emissions from space, J. Geophys. Res., 113, D16S26,
- 407 doi:10.1029/2007JD008816, 2008.
- 408
- 409 Chen, D., Wang, Y., McElroy, M. B., He, K., Yantosca, R. M., and Le Sager, P.: Regional CO
- 410 pollution and export in China simulated by the high-resolution nested-grid GEOS-Chem model,
- 411 Atmos. Chem. Phys., 9, 3825-3839, 10.5194/acp-9-3825-2009, 2009.
- 412
- 413 Cohan, D. S., Hu, Y., and Russell, A. G.: Dependence of ozone sensitivity analysis on grid
- 414 resolution, Atmos. Environ., 40, 126-135, doi:10.1016/j.atmosenv.2005.09.031, 2006.
- 415
- 416 Cooper, M., Martin, R. V., Wespes, C., Coheur, P., Clerbaux, C., and Murray, L. T.:
- 417 Tropospheric nitric acid columns from the IASI satellite instrument interpreted with a chemical
- 418 transport model: Implications for parameterizations of nitric oxide production by lightning, J.
- 419 Geophys. Res.-Atmos., 119, 10068-10079, doi:10.1002/2014JD021907, 2014.
- 420
- 421 Courant, R., Friedrichs, K., and Lewy, H.: On partial difference equations of mathematical
- 422 physics, IBM Journal of Research and Development, 11, 215-234, doi:10.1147/rd.112.0215,
- 423 1967.
- 424
- Damian, V., Sandu, M., Potra, G. R., and Carmichael: The kinetic preprocessor KPP-a software
- 426 environment for solving chemical kinetics, Comput. Chem. Eng., 26, 1567-1579,
  427 doi:10.1016/S0098-1354(02)00128-X, 2002.
- 427 d 428
- 429 Duncan, B. N., Logan, J. A., Bey, I., Megretskaia, I. A., Yantosca, R. M., Novelli, P. C., Jones,
- 430 N. B., and Rinsland, C. P.: Global budget of CO, 1988 –1997: Source estimates and validation
- 431 with a global model, J. Geophys. Res., 112, D22301, doi:10.1029/2007JD008459, 2007.
- 432

- 433 Eastham, S. D., Weisenstein, D. K., and Barrett, S. R. H.: Development and evaluation of the
- 434 unified tropospheric–stratospheric chemistry extension (UCX) for the global chemistry-transport
- 435 model GEOS-Chem, Atmos. Environ, 89, 52-63, doi:10.1016/j.atmosenv.2014.02.001, 2014.
- 436
- 437 Esler, J. G., Roelofs, G. J., Kohler, M. O., and O'Connor, F. M.: A quantitative analysis of grid-
- related systematic errors in oxidising capacity and ozone production rates in chemistry transport
- 439 models, Atmos. Chem. Phys., 4, 1781-1795, doi:10.5194/acp-4-1781-2004, 2004.
- 440
- 441 Evans, M. J., and Jacob, D. J.: Impact of new laboratory studies of N<sub>2</sub>O<sub>5</sub> hydrolysis on global
- 442 model budgets of tropospheric nitrogen oxides, ozone, and OH, Geophys. Res. Lett., 32, L09813,
- 443 doi:10.1029/2005GL022469, 2005.
- 444
- Fairlie, T. D., Jacob, D. J., and Park, R. J.: The impact of transpacific transport of mineral dust in
  the United States, Atmos. Environ., 41, 1251-1266, doi:10.1016/j.atmosenv.2006.09.048, 2007.
- 448 Fisher, J. A., Jacob, D. J., Wang, Q., Bahreini, R., Carouge, C. C., Cubison, M. J., Dibb, J. E.,
- 449 Diehl, T., Jimenez, J. L., Leibensperger, E. M., Lu, Z., Meinders, M. B. J., Pye, H. O. T., Quinn,
- 450 P. K., Sharma, S., Streets, D. G., van Donkelaar, A., and Yantosca, R. M.: Sources, distribution,
- 451 and acidity of sulfate-ammonium aerosol in the Arctic in winter-spring, Atmos. Environ., 45,
- 452 7301-7318, doi:10.1016/j.atmosenv.2011.08.030, 2011.
- 453
- 454 Fountoukis, C., Koraj, D., Denier van der Gon, H.A.C., Charalampidis, P. E., Pilinis, C., and
- 455 Pandis, S. N.: Impact of grid resolution on the predicted fine PM by a regional 3-D chemical
- 456 transport model, Atmos. Environ., 68, 24-32, doi:10.1016/j.atmosenv.2012.11.008, 2013.
  457
- Fountoukis, C., and Nenes, A.: ISORROPIA II: a computationally efficient thermodynamic
  equilibrium model for K<sup>+</sup>-Ca<sup>2+</sup>-Mg<sup>2+</sup>-NH<sub>4</sub><sup>+</sup>-Na<sup>+</sup>-SO<sub>4</sub><sup>2-</sup>-NO<sub>3</sub><sup>-</sup>-Cl<sup>-</sup>-H<sub>2</sub>O aerosols, Atmos. Chem.
  Phys., 7, 4639-4659, doi:10.5194/acp-7-4639-2007, 2007.
- 461
- Ginoux, P., Chin, M., Tegen, I., Prospero, J. M., Holben, B., Duboviki, O., and Lin, S. J.:
  Sources and distributions of dust aerosols simulated with the GOCART model, J. Geophys. Res.,
  106, 20255–20274, doi:10.1029/2000JD000053, 2001.
- 465
- 466 Heald, C. L., Collett Jr., J. L., Lee, T., Benedict, K. B., Schwandner, F. M., Li, Y., Clarisse, L.,
- 467 Hurtmans, D. R., Van Damme, M., Clerbaux, C., Coheur, P.-F., Philip, S., Martin, R. V., and
- 468 Pye, H. O. T.: Atmospheric ammonia and particulate inorganic nitrogen over the United States,
- 469 Atmos. Chem. Phys., 12, 10295-10312, doi:10.5194/acp-12-10295-2012, 2012.
- 470
- 471 Henderson, B. H., Jeffries, H. E., Kim, B. U., and Vizuete, W. G.: The influence of model
- 472 resolution on ozone in industrial volatile organic compound plumes, J. Air Waste Manage.
- 473 Assoc., 60, 1105–1117, doi:10.3155/1047-3289.60.9.1105, 2010.
- 474
- 475 Holtslag, A. A. M., and Boville, B. A.: Local Versus Nonlocal Boundary-Layer Diffusion in a
- 476 Global Climate Model, J. Clim., 6, 1825-1842, doi:10.1175/1520-
- 477 0442(1993)006<1825:LVNBLD>2.0.CO;2, 1993.

478 Horowitz, L. W., Walters, S., Mauzerall, D. L., Emmons, L. K., Rasch, P. J., Granier, C., Tie, X., 479 Lamarque, J., Schultz, M. G., Tyndall, G. S., Orlando, J. J., and Brasseur, G. P.: A global 480 simulation of tropospheric ozone and related tracers: Description and evaluation of MOZART, 481 version 2, J. Geophys. Res., 108, 4784, doi:10.1029/2002JD002853, 2003. 482 483 Hudman, R. C.; Jacob, D. J., Turquety, S., Leibensperger, E. M., Murray, L. T., Wu, S., Gilliland, 484 A. B., Avery, M., Bertram, T. H., Brune, W., Cohen, R. C., Dibb, J. E., Flocke, F. M., Fried, A., 485 Holloway, J., Neuman, J. A., Orville, R., Perring, A., Ren, X., Sachse, G. W., Singh, H. B., 486 Swanson, A., and Wooldridge, P. J.: Surface and lightning sources of nitrogen oxides over the 487 United States: Magnitudes, chemical evolution, and outflow, J. Geophys. Res., 112, D12S05, 488 doi:10.1029/2006JD007912, 2007. 489 490 Huijnen, V., Williams, J., van Weele, M., van Noije, T., Krol, M., Dentener, F., Segers, A., 491 Houweling, S., Peters, W., de Laat, J., Boersma, F., Bergamaschi, P., van Velthoven, P., Le 492 Sager, P., Eskes, H., Alkemade, F., Scheele, R., Nédélec, P., and Pätz, H. -W.: The global 493 chemistry transport model TM5: description and evaluation of the tropospheric chemistry 494 version 3.0, Geosci. Model Dev., 3, 445-473, doi:10.5194/gmd-3-445-2010, 2010. 495 496 Hundsdorfer, W., and Verwer, J.: Numerical solution of time-dependent advection-diffusion-497 reaction equations, Springer Series in Computational Mathematics, 33, 325-417, 498 doi:10.1007/978-3-662-09017-6 4, 2003. 499 500 Ito, A., Sillman, S., and Penner J. E.: Global chemical transport model study of ozone response 501 to changes in chemical kinetics and biogenic volatile organic compounds emissions due to 502 increasing temperatures: Sensitivities to isoprene nitrate chemistry and grid resolution, J. 503 Geophys. Res., 114, D09301, doi:10.1029/2008JD011254, 2009. 504 505 Jacob, D.: Heterogeneous chemistry and tropospheric ozone, Atmos. Environ., 34, 2131-2159, 506 doi:10.1016/S1352-2310(99)00462-8, 2000. 507 508 Jacobson, M. Z.: Computation of global photochemistry with SMVGEAR II, Atmos. Environ., 509 29, 2541-2546, doi:10.1016/1352-2310(95)00194-4, 1995. 510 511 Jacobson, M. Z.: Improvement of SMVGEAR II on vector and scalar machines through absolute 512 error tolerance control, Atmos. Environ., 32, 791-796, doi:10.1016/S1352-2310(97)00315-4, 513 1998. 514 515 Jacobson, M., and Turco, R. P.: SMVGEAR - A SPARSE-MATRIX, VECTORIZED GEAR 516 CODE FOR ATMOSPHERIC MODELS, Atmos. Environ., 28, 273-284, doi:10.1016/1352-517 2310(94)90102-3, 1994. 518 519 Jaegle, L., Quinn, P. K., Bates, T. S., Alexander, B., and Lin, J.-T.: Global distribution of sea salt 520 aerosols: new constraints from in situ and remote sensing observations, Atmos. Chem. Phys., 11, 521 3137-3157, doi:10.5194/acp-11-3137-2011, 2011. 522

- 523 Jang, J. C., Jeffries, H. E., and Tonnesen, S.: Sensitivity of ozone to model grid resolution II.
- 524 Detailed process analysis for ozone chemistry, Atmos. Environ., 29, 3101-3114,
- 525 doi:10.1016/1352-2310(95)00119-J, 1995.
- 526
- 527 Keller, C. A., Long, M. S., Yantosca, R. M., Da Silva, A. M., Pawson, S., and Jacob, D. J.:
- 528 HEMCO v1.0: a versatile, ESMF-compliant component for calculating emissions in atmospheric
- 529 models, Geosci. Model Dev., 7, 1409-1417, doi:10.5194/gmd-7-1409-2014, 2014.
- 530
- 531 Kraabøl, A. G., Berntsen, T. K., Sundet, J. K., and Stordal, F.: Impacts of NO<sub>x</sub> emissions from
- 532 subsonic aircraft in a global three-dimensional chemistry transport model including plume
- 533 processes, J. Geophys. Res., 107, 4655, doi:10.1029/2001JD001019, 2002.
- 534
- Li, Y., Henze, D. K., Jack, D., and Kinney, P.: The influence of air quality model resolution on
  health impact assessment for fine particulate matter and its components, Air Qual., Atmos.
  Health, 1-18, doi:10.1007/s11869-015-0321-z, 2015.
- 538
- Liang, J., and Jacobson, M. Z.: Effects of subgrid segregation on ozone production efficiency in
  a chemical model, Atmos. Environ., 34, 2975-2982, doi:10.1016/S1352-2310(99)00520-8, 2000.
- Lin, J., and McElroy, M. B.: Impacts of boundary layer mixing on pollutant vertical profiles in
  the lower troposphere: Implications to satellite remote sensing, Atmos. Environ., 44, 1726-1739,
  doi:10.1016/j.atmosenv.2010.02.009, 2010.
- 546 Lin, S., and Rood, R. B.: Multidimensional Flux-Form Semi-Lagrangian Transport Schemes,
- 547 Mon. Weather Rev., 124, 2046-2070, doi:10.1175/1520-
- 548 0493(1996)124<2046:MFFSLT>2.0.CO;2, 1996.
- 549

- Lin, S., Chao, W. C., Sud, Y. C., and Walker, G. K.: A Class of the van Leer-type Transport
  Schemes and Its Application to the Moisture Transport in a General Circulation Model, Mon.
  Weather Rev., 122, 1575-1593, doi:10.1175/1520-0493(1994)122<1575:ACOTVL>2.0.CO;2,
  1994.
- 553 554
- Liu, H. Y., Jacob, D. J., Bey, I., and Yantosca, R. M.: Constraints from Pb-210 and Be-7 on wet
- by deposition and transport in a global three-dimensional chemical tracer model driven by
- assimilated meteorological fields, J. Geophys. Res., 106, 12109-12128,
- 558 doi:10.1029/2000JD900839, 2001.
- 559
- 560 Mallet, V. and Sportisse, B.: Uncertainty in a chemistry-transport model due to physical
- 561 parameterizations and numerical approximations: An ensemble approach applied to ozone 562 modeling, J. Geophys. Res., 111, D01302, doi:10.1029/2005JD006149, 2006.
- 563
- 564 Mallet, V., Pourchet, A., Quélo, D., and Sportisse, B.: Investigation of some numerical issues in
- a chemistry-transport model: Gas-phase simulations, J. Geophys. Res., 112, D15301,
- 566 doi:10.1029/2006JD008373, 2007.
- 567

568 Mao, J., Jacob, D. J., Evans, M. J., Olson, J. R., Ren, X., Brune, W. H., St Clair, J. M., Crounse, 569 J. D., Spencer, K. M., Beaver, M. R., Wennberg, P. O., Cubison, M. J., Jimenez, J. L., Fried, A., 570 Weibring, P., Walega, J. G., Hall, S. R., Weinheimer, A. J., Cohen, R. C., Chen, G., Crawford, J. H., McNaughton, C., Clarke, A. D., Jaegle, L., Fisher, J. A., Yantosca, R. M., Le Sager, P., and 571 572 Carouge, C.: Chemistry of hydrogen oxide radicals (HO<sub>x</sub>) in the Arctic troposphere in spring, 573 Atmos. Chem. Phys., 10, 5823-5838, doi:10.5194/acp-10-5823-2010, 2010. 574 575 Mao, J., Fan, S., Jacob, D. J., and Travis, K. R.: Radical loss in the atmosphere from Cu-Fe 576 redox coupling in aerosols, Atmos. Chem. Phys., 13, 509-519, doi:10.5194/acp-13-509-2013, 577 2013. 578 579 Martin, R. V., Jacob, D. J., Yantosca, R. M., Chin, M., and Ginoux, P.: Global and regional 580 decreases in tropospheric oxidants from photochemical effects of aerosols, J. Geophys. Res., 581 108, 4097, doi:10.1029/2002JD002622, 2003. 582 583 Park, R. J., Jacob, D. J., Chin, M., and Martin, R. V.: Sources of carbonaceous aerosols over the 584 United States and implications for natural visibility, J. Geophys. Res.-Atmos., 108, 4355, 585 doi:10.1029/2002JD003190, 2003. 586 587 Park, R. J., Jacob, D. J., Field, B. D., Yantosca, R. M., and Chin, M.: Natural and transboundary 588 pollution influences on sulfate-nitrate-ammonium aerosols in the United States: Implications for 589 policy, J. Geophys. Res.-Atmos., 109, D15204, doi:10.1029/2003JD004473, 2004. 590 591 Philip, S., Martin, R.V., van Donkelaar, A., Lo, J. W.-H., Wang, Y., Chen, D., Zhang, L., 592 Kasibhatla, P. S., Wang, S. W., Zhang, Q., Lu, Z., Streets, D. G., Bittman, S., and Macdonald, D. 593 J.: Global chemical composition of ambient fine particulate matter for exposure assessment, 594 Environ. Sci. Technol., 48, 13060-13068, doi:10.1021/es502965b, 2014. 595 596 Prather, M. J.: Numerical advection by conservation of second-order moments, J. Geophys. Res., 597 91, 6671-6681, doi:10.1029/JD091iD06p06671, 1986. 598 599 Prather, M. J., Zhu, X., Strahan, S. E., Steenrod, S. D., and Rodriguez, J. M.: Quantifying errors 600 in trace species transport modeling, Proc. Natl. Acad. Sci. U. S. A., 105, 19617-19621, 601 doi:10.1073/pnas.0806541106, 2008. 602 603 Punger, E. M., and West, J. J.: The effect of grid resolution on estimates of the burden of ozone 604 and fine particulate matter on premature mortality in the USA, Air Qual., Atmos. Health, 6, 563-605 573, doi:10.1007/s11869-013-0197-8, 2013. 606 607 Pye, H. O. T., Liao, H., Wu, S., Mickley, L. J., Jacob, D. J., Henze, D. K., and Seinfeld, J. H.: 608 Effect of changes in climate and emissions on future sulfate-nitrate-ammonium aerosol levels in 609 the United States, J. Geophys. Res., 114, D01205, doi:10.1029/2008JD010701, 2009. 610 611 612

- 613 Rienecker, M. M., Suarez, M. J., Todling, R., Bacmeister, J., Takacs, L., Liu, H.-C., Gu, W.,
- 614 Sienkiewicz, M., Koster, R. D., Gelaro, R., Stajner, I., and Nielsen, J. E.: The GEOS-5 Data
- Assimilation System-Documentation of versions 5.0.1 and 5.1.0, and 5.2.0. NASA Tech. Rep.
- 616 Series on Global Modeling and Data Assimilation, NASA/TM-2008-104606, vol. 27, Goddard
- 617 Space Flight Center, Greenbelt, Maryland, USA, 92 p., 2008.
- 618
- Ridley, D. A., Heald, C. L., Pierce, J. R., and Evans, M. J.: Toward resolution-independent dust
  emissions in global models: Impacts on the seasonal and spatial distribution of dust, Geophys.
  Res. Lett., 40, 2873–2877, doi:10.1002/grl.50409, 2013.
- 622
- Rind, D., Lerner, J., Jonas, J., and McLinden, C.: Effects of resolution and model physics on
  tracer transports in the NASA Goddard Institute for Space Studies general circulation models, J.
  Geophys. Res., 112, D09315, doi:10.1029/2006JD007476, 2007.
- 626
- 627 Rotman, D. A., Atherton, C. S., Bergmann, D. J., Cameron-Smith, P. J., Chuang, C. C., Connel,
- 628 P. S., Dignon, J. E., Franz, A., Grant, K. E., Kinnison, D. E., Molenkamp, C. R., Proctor, D. D.,
- and Tannahill, J. R.: IMPACT, the LLNL 3-D global atmospheric chemical transport model for
- 630 the combined troposphere and stratosphere: Model description and analysis of ozone and other
- 631 trace gases, J. Geophys. Res., 109, D04303, doi:10.1029/2002JD003155, 2004.
- 632
- Santillana, M., Zhang, L., and Yantosca, R.: Estimating numerical errors due to operator splitting
  in global atmospheric chemistry models: Transport and chemistry, J. Comput. Phys., 305, 372386, doi:10.1016/j.jcp.2015.10.052, 2016.
- 636
- 637 Shindell, D. T., Faluvegi, G., Stevenson, D. S., Krol, M. C., Emmons, L. K., Lamarque, J.-F.,
- 638 Petron, G., Dentener, F. J., Ellingsen, K., Schultz, M. G., Wild, O., Amann, M., Atherton, C. S.,
- 639 Bergmann, D. J., Bey, I., Butler, T., Cofala, J., Collins, W. J., Derwent, R. G., Doherty, R. M.,
- 640 Drevet, J., Eskes, H. J., Fiore, A. M., Gauss, M., Hauglustaine, D. A., Horowitz, L. W., Isaksen,
- 641 I. S. A., Lawrence, M. G., Montanaro, V., Muller, J. F., Pitari, G., Prather, M. J., Pyle, J. A.,
- Rast, S., Rodriguez, J. M., Sanderson, M. G., Savage, N. H., Strahan, S. E., Sudo, K., Szopa, S.,
- 643 Unger, N., van Noije, T. P. C., and Zeng, G.: Multimodel simulations of carbon monoxide:
- 644 Comparison with observations and projected near-future changes, J. Geophys. Res., 111,
- 645 D19306, doi:10.1029/2006JD007100, 2006.
- 646
- Sillman, S., Logan, J. A., and Wofsy, S. C.: A regional scale model for ozone in the United
  States with subgrid representation of urban and power plant plumes, J. Geophys. Res., 95, 57315748, doi:10.1029/JD095iD05p05731, 1990.
- 650
- Sportisse, B.: An analysis of operator splitting techniques in the stiff case, J. Comput. Phys., 161,
  140-168, doi:10.1006/jcph.2000.6495, 2000.
- 653
- Strang, G.: On the construction and comparison of difference schemes, SIAM J. Numer. Anal., 5,
  506-517, doi:10.1137/0705041, 1968.
- 656
- 657

- Thompson, T. M., Saari, R. K., and Selin, N. E.: Air quality resolution for health impact
- assessment: influence of regional characteristics, Atmos. Chem. Phys., 14, 969-978,
- doi:10.5194/acp-14-969-2014, 2014.
- 661

van Donkelaar, A., Zhang, L., Chen, D., Martin, R. V., Pasch, A. N., Szykman, J. J., and Wang,
Y. X.: Improving the accuracy of daily satellite-derived ground-level fine aerosol concentration
estimates for North America, Environ. Sci. Technol., 46, 11971-11978, doi:10.1021/es3025319,
2012.

- 666
- Valin, L. C., Russell, A. R., Hudman, R. C., and Cohen, R. C.: Effects of model resolution on the
  interpretation of satellite NO<sub>2</sub> observations, Atmos. Chem. Phys., 11, 11647-11655,
  doi:10.5194/acp-11-11647-2011, 2011.
- 670
- Vinken, G. C. M., Boersma, K. F., Jacob, D. J., and Meijer, E. W.: Accounting for non-linear
- chemistry of ship plumes in the GEOS-Chem global chemistry transport model, Atmos. Chem.
  Phys., 11, 11707-11722, doi:10.5194/acp-11-11707-2011, 2011.
- 674
- Wang, Q., Jacob, D. J., Fisher, J. A., Mao, J., Leibensperger, E. M., Carouge, C. C., Le Sager, P.,
- Kondo, Y., Jimenez, J. L., Cubison, M. J., and Doherty, S. J.: Sources of carbonaceous aerosols
- and deposited black carbon in the Arctic in winter-spring: implications for radiative forcing,
- 678 Atmos. Chem. Phys., 11, 12453-12473, doi:10.5194/acp-11-12453-2011, 2011.
- 679
- Wang, H., Jacob, D. J., Le Sager, P., Streets, D. G., Park, R. J., Gilliland, A. B., and van
- 681 Donkelaar, A.: Surface ozone background in the United States: Canadian and Mexican pollution
- 682 influences, Atmos. Environ., 43, 1310–1319, doi:10.1016/j.atmosenv.2008.11.036, 2009.
- 683
- Wang, Y., Jacob, D. J., and Logan, J. A.: Global simulation of tropospheric O<sub>3</sub>-NO<sub>x</sub>-hydrocarbon
  chemistry: 1. Model formulation, J. Geophys. Res., 103, 10713-10725, doi:10.1029/98JD00158,
  1998.
- 687
- Wang, Y. X., McElroy, M. B., Jacob D. J., and Yantosca, R. M.: A nested grid formulation for
- 689 chemical transport over Asia: Applications to CO, J. Geophys. Res., 109,
- 690 doi:10.1029/2004JD005237, 2004.
- 691
- Wesely, M. L.: Parameterization of surface resistances to gaseous dry deposition in regionalscale numerical models, Atmos. Environ., 23, 1293-1304, doi:10.1016/0004-6981(89)90153-4,
  1989.
- 695
- Wild, O., and Prather, M. J.: Global tropospheric ozone modeling: Quantifying errors due to grid
  resolution, J. Geophys. Res., 111, D11305, doi:10.1029/2005JD006605, 2006.
- 698
- Wu, S., Mickley, L. J., Jacob, D. J., Logan, J. A., Yantosca, R. M. and Rind, D.: Why are there
- large differences between models in global budgets of tropospheric ozone?, J. Geophys. Res.,
   112 D05202 doi:10.1020/2006/D007801.2007
- 701 112, D05302, doi:10.1029/2006JD007801, 2007. 702
  - 24

Yan, Y.-Y., Lin, J.-T., Kuang, Y., Yang, D., and Zhang, L.: Tropospheric carbon monoxide over
 the Pacific during HIPPO: two-way coupled simulation of GEOS-Chem and its multiple nested

- 705 models, Atmos. Chem. Phys., 14, 12649-12663, doi:10.5194/acp-14-12649-2014, 2014.

Yu, K., Jacob, D. J., Fisher, J. A., Kim, P. S., Marais, E. A., Miller, C. C., Travis, K. R., Zhu, L.,
Yantosca, R. M., Sulprizio, M. P., Cohen, R. C., Dibb, J. E., Fried, A., Mikoviny, T., Ryerson, T.
B., Wennberg, P. O., and Wisthaler, A.: Sensitivity to grid resolution in the ability of a chemical
transport model to simulate observed oxidant chemistry under high-isoprene conditions, Atmos.

711 Chem. Phys. Discuss., doi:10.5194/acp-2015-980, in review, 2016.

713 Zender, C. S., Bian, H., and Newman, D.: The mineral dust entrainment and deposition (DEAD)

- model: description and 1990s dust climatology, J. Geophys. Res., 108, 4416,
- 715 doi:10.1029/2002JD002775, 2003.

- scheme for an atmospheric aerosol module, Atmos. Environ., 35, 549-560, doi:10.1016/S13522310(00)00326-5, 2001.

721 Zhang, L., Jacob, D. J., Downey, N. V., Wood, D. A., Blewitt, D., Carouge, C. C., van

722 Donkelaar, A., Jones, D. B. A., Murray, L. T., and Wang, Y.: Improved estimate of the policy-

relevant background ozone in the United States using the GEOS-Chem global model with  $1/2^{\circ} \times$ 

- 2/3° horizontal resolution over North America, Atmos. Environ., 45, 6769-6776,
- 725 doi:10.1016/j.atmosenv.2011.07.054, 2011.

<sup>717</sup> Zhang, L. M., Gong, S. L., Padro, J. and Barrie, L.: A size-segregated particle dry deposition

747 Table 1: Comparison of mean\* relative simulation error versus horizontal resolution, with "truth"

# 748 defined at $2^{\circ} x 2.5^{\circ}$ horizontal resolution

	Species	Mean relative simulation error (unitless)	
		4° x 5° resolution	2° x 2.5° resolution
	Nitrogen oxides	2.1	0.092
	Secondary inorganic aerosols	1.0	0.14
	Ozone	0.17	0.004
	Carbon monoxide	0.36	0.005
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751	* Mean taken for operator durations $\leq$	30 min.	
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Figure 1: CPU time for GEOS-Chem simulations with various operator durations at three horizontal resolutions. Global simulations are at  $4^{\circ} \times 5^{\circ}$  (top) and  $2^{\circ} \times 2.5^{\circ}$  (middle) resolutions. The bottom panel contains results for the average of two nested regions North America and East Asia at  $0.5^{\circ} \times 0.667^{\circ}$  resolution. Colored lines represent the CPU time for simulating transport (red) and chemical (blue) operators, and the sum of the two (green). Error bars represent standard error over five simulations. Simulations are represented in the abscissa as CccTtt with chemical operator duration, C = cc minutes, and transport operator duration, T = tt minutes.

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775 Figure 2a: Sensitivity of simulated species to the duration of chemical and transport operators. The 776 left column contains monthly mean ground-level concentrations simulated with the shortest 777 operator duration considered (C10T05) at 2° x 2.5° horizontal resolution. Other columns contain 778 the absolute differences from doubling the transport operator duration to C10T10 (middle), and 779 doubling the chemical operator duration to C20T05 (right). Each row from top to bottom 780 represents carbon monoxide (CO), nitrogen oxides (NO<sub>x</sub>), hydroxyl radical (OH), and the 781 production of ozone (P[O3]). Simulations are represented as CccTtt with chemical operator 782 duration, C = cc minutes, and transport operator duration, T = tt minutes.



785 Figure 2b: As described in Fig. 2a, but each row from top to bottom represents ozone  $(O_3)$ , sulfur

<sup>786</sup> dioxide (SO<sub>2</sub>), sulfate (SO<sub>4</sub><sup> $2^-$ </sup>), and nitrate (NO<sub>3</sub><sup>-</sup>).



794 Figure 3: Sensitivity of simulated species to changes in operator duration (C20T10 to C10T05) at 795 two different horizontal resolutions over North America (global 4° x 5°, and nested 0.5° x 0.67° 796 simulations). The upper two rows contain monthly mean ground-level concentrations simulated 797 with the C20T10 operator duration for 4° x 5° (top row) and 0.5° x 0.67° (second row) resolutions. The two lower rows contain the monthly mean differences (C20T10 minus C10T05) for 4° x 5° 798 799 (third row) and 0.5° x 0.67° (bottom row) resolutions. Each column from left to right represents 800 nitrogen oxides (NO<sub>x</sub>), secondary inorganic aerosols (SIA), ozone (O<sub>3</sub>), and carbon monoxide 801 (CO).



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Figure 4: Relative simulation error of different species ( $E_{sim}^s$ , eq. 3) with various operator durations at 2° x 2.5° horizontal resolution. Colored lines and dots represent the relative simulation error for nitrogen oxides (NO<sub>x</sub>; red), secondary inorganic aerosols (SIA; blue), ozone (O<sub>3</sub>; green), and carbon monoxide (CO; magenta). Simulations are represented in the abscissa as CccTtt with chemical operator duration, C = cc minutes, and transport operator duration, T = tt minutes.

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815 Figure 5: Effect on simulated species of changing from the GEOS-Chem traditional operator 816 durations (C30T15) to the shortest operator durations considered (C10T05). The top row contains 817 monthly mean ground-level concentrations simulated with the C30T15 operator duration at 2° x 818 2.5° horizontal resolution. The next two rows contain the monthly mean differences (C30T15 819 minus C10T05) for absolute (second row) and relative (third row) differences. The two lowest 820 rows contain the maximum differences (C30T15 minus C10T05) for absolute (fourth row) and 821 relative (bottom row) differences. Each column from left to right represents nitrogen oxides (NO<sub>x</sub>), 822 secondary inorganic aerosols (SIA), ozone (O<sub>3</sub>), and carbon monoxide (CO).



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Figure 6: CPU-time adjusted Composite Normalized Error (CNE, eq. 4) for GEOS-Chem simulations with various horizontal resolutions and operator durations. Colored lines and dots represent the CNE for the global simulations at  $4^{\circ} \times 5^{\circ}$  (red) and  $2^{\circ} \times 2.5^{\circ}$  (blue), and the nested simulations at  $0.5^{\circ} \times 0.667^{\circ}$  (green) horizontal resolutions. Error bars represent standard error in CPU time. Simulations are represented in the abscissa as CccTtt with chemical operator duration, C = cc minutes, and transport operator duration, T = tt minutes.