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mizuRoute version 1: a river network routing tool for a continental domain water resources applications

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Abstract

This paper describes the first version of a stand-alone runoff routing tool, mizuRoute, which post-processes runoff outputs from any distributed hydrologic model or land surface model to produce spatially distributed streamflow at various spatial scales from headwater basins to continental-wide river systems. The tool can utilize both traditional grid-based river network and vector-based river network data, which includes river segment lines and the associated drainage basin polygons. Streamflow estimates at any desired location in the river network can be easily extracted from the output of mizuRoute. The routing process is simulated as two separate steps. The first is hillslope routing, which uses a gamma distribution to construct a unit-hydrograph that represents the transport of runoff from a hillslope to a catchment outlet. The second step is river channel routing, which is performed with one of two routing scheme options: (1) a kinematic wave tracking (KWT) routing procedure; and (2) an impulse response function–unit hydrograph (IRF-UH) routing procedure. The mizuRoute system also includes tools to pre-process spatial river network data. This paper demonstrates mizuRoute's capabilities with spatially distributed streamflow simulations based on river networks from the United States Geological Survey (USGS) Geospatial Fabric (GF) dataset, which contains over 54 000 river segments across the contiguous United States (CONUS). A brief analysis of model parameter sensitivity is also provided. The mizuRoute tool can assist model-based water resources assessments including studies of the impacts of climate change on streamflow.

1 Introduction

The routing tool described in this paper post-processes runoff outputs from macro-scale hydrologic models or land surface models (hereafter we use "hydrologic model" to refer to both types of model) to estimate spatially distributed streamflow along the river network. The river routing tool is named mizuRoute ("*mizu*" means "*water*" in

Japanese). The motivation for mizuRoute's development was to enable continental domain evaluations of hydrologic simulations for water resources assessments, such as studies of the impacts of climate change on streamflow. The mizuRoute tool is suitable for processing of ensembles of multi-decadal runoff outputs because the tool is standalone and easily applied in a parallel mode. The mizuRoute tool is also designed to output streamflow estimates at all river segments in the river network across the domain of interest at each time step, facilitating the further spatial and temporal analysis of the estimated streamflow.

The paper proceeds as follows. Section 2 reviews existing river routing models. Section 3 describes the hillslope and river routing schemes used in mizuRoute. Section 4 provides an overview of the workflow of mizuRoute from preprocessing hydrologic model output to simulating streamflow in the river network. Section 5 demonstrates streamflow simulations in river systems over the contiguous United States. Finally, a summary and future work are discussed in Sect. 6.

2 Existing river routing models

The water resources and Earth System Modeling communities have developed a wide spectrum of river routing schemes of varying complexity (Clark et al., 2015). For example, the US Army Corps of Engineers (USACE) has developed a stand-alone river modeling system called Hydrologic Engineering Center-River Analysis System (HEC-RAS; Brunner, 2001). HEC-RAS offers various hydraulic routing schemes, ranging from simple uniform flow to one-dimensional (1-D) Saint-Venant equations for unsteady flow. HEC-RAS has been popular among civil engineers for river channel design and floodplain analysis where surveyed river geometry and physical channel properties are available. At the continental to global scale, unit-hydrograph approaches have been used (e.g., Nijssen et al., 1997; Lohmann et al., 1998; Goteti et al., 2008; Zaitchik et al., 2010; Xia et al., 2012), though more recent, large-scale river models use fully dynamic flow equations (e.g., Miguez-Macho and Fan, 2012; Paiva et al., 2013;

Clark et al., 2015), simplified Saint-Venant equations such as the kinematic wave or diffusive wave equation (e.g., Arora and Boer, 1999; Lucas-Picher et al., 2003; Koren et al., 2004; Yamazaki et al., 2011, 2013; Li et al., 2013; Gochis, 2015; Yucel et al., 2015) or non-dynamical hydrologic routing methods such as Muskingum routing (e.g., David et al., 2011). Despite their computational cost, dynamic or diffusive wave models are attractive for relatively flat floodplain regions such as along the Amazon River where backwater effects on the flood wave are significant (Paiva et al., 2011; Yamazaki et al., 2011; Miguez-Macho and Fan, 2012). At the other end of the spectrum, simpler, non-dynamic routing schemes, such as the unit hydrograph approach, estimate the flood wave delay and attenuation, but no other streamflow variables such as flow velocity and flow depth.

One of the key issues for large scale river routing, besides the choice of the routing scheme, is the degree of abstraction in the representation of the river network (Fig. 1). A vector-based representation of the river network refers to a collection of hydrologic response units (HRUs) that are delineated based on topography or catchment boundary. River segments in the vector-based river network, represented by lines, meander through HRUs and connect upstream with downstream HRUs. On the other hand, in the grid-based river network, the HRU is represented by a grid box and river segments connect neighboring grid boxes based on the flow directions. Vector-based river networks are better than coarser resolution (e.g. > 1 km) gridded river networks at preserving fine-scale features of the river system such as tortuosity, therefore representing more accurate sub-catchment areas and river segment lengths.

For large scale applications, many studies have developed and evaluated methods to upscale fine resolution flow direction grids (~ 1 km or less) to a coarser resolution (~ 10 km or more) to match hydrologic model resolution and/or reduce the cost of routing computations (e.g., O'Donnell et al., 1999; Fekete et al., 2001; Olivera et al., 2002; Reed, 2003; Davies and Bell, 2009; Wu et al., 2011, 2012). Earlier work (e.g., O'Donnell et al., 1999; Fekete et al., 2001; Olivera et al., 2002) focused on preserving the accuracy of the flow direction at the coarser resolution and therefore on an accurate representation of the drainage area. Newer upscaling methods are designed to also preserve fine-scale flow path length (e.g., Yamazaki et al., 2009; Wu et al., 2011). More recent river routing models have also begun to employ vector-based river networks (Goteti et al., 2008; David et al., 2011; Paiva et al., 2011, 2013; Lehner and Grill, 2013; Yamazaki et al., 2013).

3 Runoff routing in mizuRoute

The runoff routing in mizuRoute provides more flexibility in continental domain routing applications. The mizuRoute framework enables model flexibility in two ways: first, mizuRoute can be used to simulate streamflow for both grid- and vector-based river networks. Given either type of river network data, mizuRoute offers an option to route flow along all the river segments in the river network data or route runoff at an outlet segment specified by a user. With the latter option, routing computation is performed only in the upstream river network of the specified outlet, which reduces the computational cost. Second, the modular structure of the mizuRoute framework offers the flexibility to configure multiple routing schemes. The current version of mizuRoute includes two different type of river routing schemes: (1) kinematic wave tracking (KWT) routing and (2) impulse response function–unit hydrograph (IRF-UH) routing, mimicking the Lohmann et al. (1996) model. This flexibility offers new capabilities not present in existing routing models. One capability is to provide an opportunity to explore routing model uncertainties originating from the representation of the river system and routing scheme differences (equations and parameters) separately.

The mizuRoute tool uses a two-step process to route basin runoff. First, basin runoff is routed from each hillslope to the river channel using a gamma-distribution-based unit-hydrograph. This allows the representation of ephemeral channels or channels too small to be included in the river network. Second, using one of the two channel routing schemes, the delayed flow from each HRU is routed to downstream river segments along the river network. The routing time step is the same as runoff output from the hydrologic model, typically an hourly or daily time step. The following sub-sections provide descriptions of the two routing steps.

3.1 Hillslope routing

Hillslope routing accounts for the time of concentration (Tc) of a local catchment (i.e., an HRU) to estimate temporally delayed runoff (or discharge) at the outlet of the HRU from runoff computed by a hydrologic model.

For hillslope routing mizuRoute uses a simple two-parameter Gamma distribution as a unit-hydrograph to route instantaneous runoff from a hydrologic model to a HRU outlet. The Gamma distribution is expressed as:

$$\gamma(t:a,\theta) = \frac{1}{\Gamma(a)\,\theta^a} t^{a-1} e^{-\frac{t}{\theta}} \tag{1}$$

where t is time [T], a is a shape parameter [–] ($a > 0$), and θ is a time-scale parameter [T]. Both the shape and time scale parameters determine the peak time (mean of the distribution: $a\theta$) and flashiness (variance of the distribution: $a\theta^2$) of the unit-hydrograph and depend on the physical catchment characteristics. Convolution of the gamma distribution with the runoff depth series is used to compute the fraction of runoff at the current time which is discharged to its corresponding river segment at each future time as follows:

$$q(t) = \int_0^{\mathrm{tmax}} \gamma(s:a,\theta)\cdot R(t-s)\,\mathrm{d}s \tag{2}$$

where q is delayed runoff or discharge [$\mathrm{L}^3\,\mathrm{T}^{-1}$] at time step t [T], R is HRU total runoff depth [$\mathrm{L}^3\,\mathrm{T}^{-1}$] from hydrologic model, and tmax is the maximum time length for the gamma distribution [T].

3.2 River channel routing

Two different river channel routing schemes are implemented in mizuRoute: (1) the kinematic wave tracking (KWT) routing procedure; and (2) the impulse response function–unit hydrograph (IRF-UH) routing procedure. Both schemes are based on the 1-D Saint-Venant equations that describe flood wave propagation through a river channel. The one-dimensional conservation equations for continuity (Eq. 3) and momentum (Eq. 4) are

$$\frac{\partial q}{\partial x} + \frac{\partial A}{\partial t} = 0 \tag{3}$$

$$\frac{\partial v}{\partial t} + v\frac{\partial v}{\partial x} + g\frac{\partial y}{\partial x} - g(S_0 - S_\mathrm{f}) = 0 \tag{4}$$

where q is discharge [$\mathrm{L}^3\,\mathrm{T}^{-1}$] at time step t [T] and location x [L] in a river network, A is cross-sectional flow area [L^2], v is velocity [$\mathrm{L}\,\mathrm{T}^{-1}$], y is depth of flow [L], S_0 is channel slope [–], S_f is friction slope [–], and g is gravitational constant [$\mathrm{L}\,\mathrm{T}^{-2}$]. The continuity equation (Eq. 3) assumes that no lateral flow is added to a channel segment. The following sub-sections describe the two routing schemes.

3.2.1 Kinematic wave tracking (KWT)

In contrast with several other kinematic routing models that solve a kinematic wave equation with the numerical schemes (e.g., Arora and Boer, 1999; Lucas-Picher et al., 2003; Koren et al., 2004), the KWT method computes a wave speed or a celerity for the runoff (or discharge) that enters an individual stream segment from the corresponding HRU at each time step using kinematic approximation (Goring, 1994; Clark et al., 2008). This runoff is tracked along the river network with a wave based on a travel time (the celerity divided by the segment length) once in the river segment. Note that the wave celerity differs from the flow velocity, as the wave typically moves faster than water mass (McDonnell and Beven, 2014).

In the kinematic wave approximation with the assumption that the channel is rectangular and hydraulically wide (channel width $\gg y$), the wave celerity C [L T^{-1}] is a function of channel width w [L], Manning's coefficient n [–], channel slope S_0 [–] and discharge q [L^3 T^{-1}]. Further details are provided in Appendix A. Among the four variables, the channel slope S_0 is provided in the river network data and discharge is computed with hillslope routing for the headwater basin, or/and updated via routing from the upstream segment. The other two variables, Manning's coefficient n and river width w, are much more difficult to measure or estimate. The river width is determined with the following width-drainage area relationship (Booker, 2010):

$$w = W_a \cdot A_{\mathrm{ups}}^b \tag{5}$$

where W_a is a width factor [–], A_{ups} is the total upstream basin area [L^2] and b is an empirical exponent equal to 0.5. The width factor W_a and the Manning's coefficient n are treated as model parameters as shown in Table 1.

The KWT routing starts with ordering all the segments in the processing sequence from upstream to the downstream segments. The KWT routing is performed at each segment in the processing order at each time step. The procedures of the KWT routing method are detailed as follows:

1. The first routine obtains the information on the waves that reside in the segment at a given time step: the waves routed from the upstream segments, the wave that remains in this current segment form the previous time step, and the wave generated from the runoff from local HRUs during the current time step. Three state variables of the waves are kept in the memory: discharge, time at which the wave enters the segment, and time at which the wave is expected to exit the segment (assign the missing value to the waves routed from the upstream segment, and computed in step 3). At the first time step, only wave from local HRUs exists. Figure 2a visualizes the discharge of waves that reside in the 16 segments (16 river segments are shown in the inserted map) at the beginning of 5 time steps against the wave locations in the segments.

2. The second routine (wave removal routine) reduces the memory usage as well as the processing time for the wave routing (the next step). The number of the waves in the segment is limited to a predefined number (20 by default). In Fig. 2, the threshold of the number of wave in the segment is set to 100. To determine which waves can be removed, first, difference between the discharge of the wave and linearly interpolated discharge values between its two neighboring waves is computed for all the waves, and then the wave that produces the least difference (from the interpolated discharge) is removed so that loss of wave mass is minimized. This process is repeated until the number of waves become lower than the threshold.

3. The third routine performs the wave routing over a given river segment. In the routing routine, the celerity of each wave in the segment is computed with Eq. (A6), and then the time at which each wave is expected to exit the river segment is updated. If the exit time occurs before the end of the time step, the wave is propagated to the downstream segment and flagged as "exited". The exit time then becomes the time the wave entered the downstream segment. Otherwise, the wave is flagged as "not-exited", and remains in the current segment. Figure 2b shows the discharge of the waves against the exit times of the corresponding waves at segments 4 and 13. As a reference, the end of each time step is shown as a vertical line. In Fig. 2b, the waves situated before the end of time step exits the segments at a given time step. The routing routine checks for (and corrects) the special case of a kinematic shock. A kinematic shock is a sudden rise in the flow depth, thus an increase in the discharge at a fixed location, and occurs when a faster-moving wave successively overtakes multiple slower-waves to build a steep wave front. It occurs in models due to the kinematic approximation; in reality, diffusion would act to reduce the steepness of the shock. Two neighboring waves are evaluated to check if a slower wave is overtaken by a faster wave before the waves exit the river segment. If this occurs, those two waves are merged into

one, and the celerity of the merged wave is updated with the following equation;

$$C_{\text{merge}} = \frac{\Delta q}{\Delta A} = \frac{\Delta q}{w \Delta y} \tag{6}$$

where C_{merge} is the merged wave celerity [$\mathrm{L\,T^{-1}}$], Δq and ΔA are differences in discharge [$\mathrm{L^3 T^{-1}}$] and cross-sectional flow area [$\mathrm{L^2 T^{-1}}$], respectively, between slower and faster waves. Note that Eq. (3) is the mathematical definition of the wave celerity. Since we assume the rectangular channel whose width is constant for each segment, the merged celerity C_{merge} is a function of the flow depth y, which is computed with Eq. (A3).

4. Finally, the time step averaged discharge (streamflow) is computed by temporal integration of discharge of all the waves that exit the segment during the time step. Temporal integral of wave discharge is visualized in Fig. 2b as the area enclosed by the discharge curve formed by all the exiting waves between the beginning and end of time step.

3.2.2 Impulse response function-unit hydrograph (IRF-UH)

The IRF-UH method mimics the river routing model of Lohmann et al. (1996), which has been used to route flows from gridded land surface models such as the Variable Infiltration Capacity model (VIC; Liang et al., 1996). The only difference between the current tool and the Lohmann routing tool is the way in which the river network is defined. The Lohmann routing model is designed as a grid-based model as shown in Fig. 1 to ease the coupling with grid-based land surface models. In mizuRoute, the same IRF-UH method can be used either on a vector- or grid-based river network. Here, the descriptions of IRF-UH are given briefly.

The mathematical developments of IRF-UH are based on one-dimensional diffusive wave equation derived from the 1-D Saint-Venant equations (Eqs. 4 and 5):

$$\frac{\partial q}{\partial t} = D\frac{\partial^2 q}{\partial x^2} - C\frac{\partial q}{\partial x} \tag{7}$$

where parameters C and D are wave celerity [$\mathrm{L\,T^{-1}}$] and diffusivity [$\mathrm{L^2\,T^{-1}}$], respectively. The complete derivation from Eqs. (4) and (5) to Eq. (7) is given in Appendix B.

Equation (4) can be solved using convolution integrals

$$q = \int_0^t U(t-s)\,h(x,s)\,\mathrm{d}s \tag{8}$$

where

$$h(x,t) = \frac{x}{2t\sqrt{\pi D t}}\exp\left(-\frac{(Ct-x)^2}{4Dt}\right) \tag{9}$$

and $U(t-s)$ is a unit depth of runoff generated at time $t-s$. This solution is a mathematical representation of the impulse response function (IRF) used in unit hydrograph theory. Wave celerity C and diffusivity D are treated as input parameters for this tool (Table 1), and ideally they can be estimated from observations of discharge and channel geometries at gauge locations.

4 mizuRoute workflow

Overall workflow of mizuRoute is illustrated in Fig. 3. There are two main separate data preprocesses before executing the main executable of mizuRoute, route_runoff.exe.

First, if the hydrologic model simulations are performed with spatial discretization (i.e., model HRU) different than the HRU used in the river network data, it is necessary to map the runoff output from the hydrologic models to the river network HRUs. This process is done by taking the area-weighted runoff of the intersecting hydrologic model HRUs. We developed the python scripts to identify the intersected hydrologic model HRUs for each river network HRU and their fractional areas to the river network HRU area to assist with this process.

The second data pre-processing step is augmentation of the river network dataset. Typical topological information in this dataset is the immediate downstream segment for each segment. While a river network can be fully defined based on information about the immediate downstream segment, the river routing schemes in mizuRoute require identification of all the upstream river segments. For this purpose, we have developed a program (process_river_topology.exe) that identifies all the upstream segments for each segment in the river network data based on the information on immediate downstream segment. This identification of upstream segments only has to be done once for each unique river network dataset. Therefore, the program (i.e., process_river_topology.exe) can be used as a preprocessor, which improves the efficiency of the main routing tool, especially when the routing is performed for multiple hydrologic model outputs for a large river system. In addition to the identification of all upstream segments, the topology program identifies upstream HRUs, upstream areas (cumulative area of all the upstream HRUs), total upstream distance from each segment to all the upstream segments, etc.

5 CONUS-wide mizuRoute simulations

This section demonstrates the capabilities of mizuRoute using the United States Geological Survey (USGS) Geospatial Fabric (GF) vector-based river network (Viger, 2014; http://wwwbrr.cr.usgs.gov/projects/SW_MoWS/GeospatialFabric.html), applied over the contiguous United states (CONUS). The river routing scheme uses both KWT and IRF-UH. In addition, sensitivity of the streamflow estimates to the river routing parameters is examined at selected locations. Different routing model choices (routing scheme and parameters) will differently affect the attenuation of runoff (i.e., the magnitude of peak and rate of rising and recession limbs) and the timing of the peak flow. Note that accuracy of the routed flow is not discussed because it depends largely on the accuracy of runoff estimates from the hydrologic model.

5.1 The Geospatial Fabric network topology

The GF dataset was developed primarily to facilitate CONUS-wide hydrologic modeling with the USGS Precipitation Runoff Modeling System (PRMS; Leavesley and Stannard, 1995). To reduce the computational burden of the hydrologic simulations, the GF dataset is generated by aggregating fine-scale river segments and corresponding catchments or HRUs from the first version of National Hydrography Dataset Plus (NHDPlus v1; HorizonSystemsCorporation 2010), while still representing small catchments (equivalent in area to 12 digit Hydrologic Unit Code $\sim 100\,\mathrm{km}^2$ or smaller basin). The GF dataset includes line and polygon geometries representing river segments and their catchments, respectively, along with their attribute information including the connectivity between segments (topological information) and their physical attributes such as channel length, area of the catchment. Table 2 lists the variables of river network vector data necessary for the mizuRoute. The GF dataset (both geometry and attribute information) is stored in Environmental System Research Institute (ESRI) Geodatabase Feature Classes and the topological and physical data (Table 2) in the attribute table is converted to NetCDF format to start with the augmentation of rive network topology (Fig. 3). The GF dataset include 54 929 river segments and 106 973 catchment HRUs (including the right and left bank of each segment). Figure 4 displays distribution of river segments in the GF vector data. The upstream area of each river segment shown in Fig. 4 is computed based on drainage based HRU (not shown in the Fig. 4) provided in the GF dataset. Although this paper illustrates runoff routing using GF, the mizuRoute tool can work with any other river network data if it includes correspondence between HRU and segment as well as segment-to-segment topology information.

5.2 Model setup

We routed the daily runoff simulations archived by Reclamation (2014) as part of their project "Downscaled CMIP3 and CMIP5 Climate and Hydrology Projections" (http://gdo-dcp.ucllnl.org/downscaled_cmip_projections/dcpInterface.html). In this project, the VIC model was forced by the spatially downscaled temperature and precipitation outputs at 1/8° resolution from 97 global climate model outputs from 1950 through 2099. Additionally, historical runoff simulations were produced at 1/8° resolution by the VIC model forced by meteorological forcings from Maurer et al. (2002) from 1950 through 1999 (Maurer et al. data is referred to as M02). The details of the Coupled Model Intercomparison Project Phase 5 (CMIP5) are described by Taylor et al. (2011).

First, to use the GF vector-based river network, the 1/8° gridded runoff was mapped to each GF HRU by taking areal weighted average of the intersecting area between grid boxes and the GF HRUs.

The routing parameters for each scheme (see Table 1) need to be predetermined. The channel parameters included in the KWT routing method (Manning's coefficient, n, and river width, w) can be determined by a survey of river channel geometry and river bed condition if the spatial scale of the model domain is very small, but this is usually infeasible for large spatial domains such as the entire CONUS used here. For the IRF-UH method, the determination of celerity and diffusivity with Eq. (B8) requires information on flow and channel geometry, so for simplicity we follow Lohmann et al. (1996) and treat celerity and diffusivity as parameters. For both schemes, parameter estimation methods need to be developed to determine appropriate values for large-scale applications. For this simulation, the parameter values are determined arbitrarily, with the objective to demonstrate the capabilities of mizuRoute to produce spatially distributed streamflow, not to attain the most accurate simulation.

The mizuRoute tool outputs the time series of the streamflow estimates at all the river segments in the river network, and modeled streamflow for the point of interest (e.g., streamflow gauge location) can be extracted from the NetCDF output file with the ID of the river segment (i.e., seg_id) where the point of interest is located.

5.3 Spatially distributed streamflow in the river network

Here we show spatially distributed streamflow estimates using the VIC simulated runoff forced with M02 meteorological data. Figure 5 shows daily mean streamflow estimated with KWT and IFR-UH routing methods for 15 June 1986 as an example. As shown in Fig. 5, both routing schemes produce qualitatively the same spatial pattern of the daily streamflow. From the spatially distributed streamflow time series, point streamflow time series are easily extracted as illustrated in Fig. 6. Figure 6 shows daily streamflow from 1 January 1995 to 31 December 1999 at three locations: (A) Snake River below Ice Habor Dam, (B) Colorado River at Lees Ferry, (C) Apalachicola River near Blountstown. Temporal patterns of flow simulations with the two river routing schemes are very similar, but the day-to-day differences in estimated streamflow due to the different routing choices become visible.

Another demonstration of mizuRoute's capability is to produce an ensemble of projected streamflow estimates from the runoff simulations using CMIP5 data. Figure 7 shows the monthly mean of 28 projected streamflow estimates (using CMIP5 RCP 8.5 scenario) at the three locations over three periods: (P1) from 2010 to 2039, (P2) from 2040 to 2069, and (P3) from 2070 to 2099. In this example, the results from the KWT scheme are shown in the Fig. 7. The interpretation of the climate changes impact on the streamflow is not discussed here and complete analyses are left for future investigation.

5.4 Sensitivity of streamflow estimates to river routing parameters

Analysis of the sensitivity of simulated hydrographs to channel routing parameters (Table 1) is performed to examine the effect of parameter values on the streamflow simulations. In this paper, brief sensitivity simulations were performed using VIC simulated runoff with M02 data and using different river routing parameter values (two parameters for each scheme). We carried out the parameter sensitivity analysis at the three locations in Fig. 6, but found the characteristics of the parameter sensitivity are the same. Therefore, we present the results for Colorado River at Lees Ferry. Figure 8 shows effect of width factor W_a in Eq. (6) (top panels) and Manning coefficient n (bottom panels) for the KWT scheme. As expected, wider channel width (with larger W_a value) produces later hydrograph shifts because larger flow area produces slower velocity to conserve the amount of discharge. This effect is enhanced with larger manning coefficient n due to more friction slowing down water flow with larger Manning coefficient n. A similar effect is seen for sensitivity of simulated hydrographs to the Manning coefficient n (bottom panel of Fig. 8).

Figure 9 shows the sensitivity of a simulated hydrograph from 1 October 1990 to 30 September 1991 to the two IRF-UH parameters at Colorado River at Lees Ferry (top panel for sensitivity to celerity C and bottom panel for sensitivity to diffusivity D). Interestingly, the effect of diffusivity D is small while celerity C affects timing of hydrograph peak. This is because celerity C directly changes peak timing without attenuation of IRF, while diffusivity D has little influence on peak timing of IRF although it changes the degree of flashiness (Eq. 12). Due to the low sensitivity of the hydrograph to diffusivity D, the degree of hydrograph sensitivity to celerity C is consistent across different diffusivity values (bottom panel of Fig. 8).

6 Summary and discussion

This paper presents mizuRoute (version 1.0), a river network routing tool that post-processes runoff outputs from any hydrologic or land surface model. We demonstrated the capability of mizuRoute to produce spatially distributed streamflow on a vector-based river network using the USGS GF river network over the CONUS. The streamflow time series are easily extracted at any locations in the network, facilitating hydrologic modeling evaluation, and other hydrologic assessments. The tool is independent of the hydrologic simulations, making it possible to produce ensembles of streamflow estimations from multiple hydrologic models. As an example of a practical application of mizuRoute, an ensemble of streamflow projections was produced at USGS gauge points on the river systems across the CONUS from 97 runoff simulations from Downscaled CMIP5 Climate and Hydrology Projections (Reclamation 2014). Section 5.3 shows some of the streamflow simulations based on the runoff generated with VIC forced by CMIP5 data.

Based on the simulations presented in the Sect. 5.4, the routing parameters can affect the simulated hydrograph especially for the KWT method. Though more detailed investigations of those effects need to be performed to fully understand the routing model behaviors, the parameter sensitivity is substantial. More sophisticated methods to estimate routing model parameters need to be developed. River physical parameters are difficult to obtain in a consistent way at the continental scale, but recent developments of the retrieval algorithms for river physical properties (channel width, slope etc.) with remote sensing data are promising (e.g., Pavelsky and Smith, 2008; Fisher et al., 2013; Allen and Pavelsky, 2015), and we expect to see advances in capabilities to estimate the hydraulic geometry of rivers over the coming years (Clark et al., 2015).

Appendix A: Derivation of wave celerity equation used in KWT

The kinematic wave approximation of the full Saint-Venant equations (Eqs. 3 and 4) uses the continuity equation combined with a simplified momentum equation. The simplified momentum equation is based on the assumption that the friction slope is equal to the channel slope and that flow is steady and uniform. Under this assumption, Eq. (4) is reduced to $S_0 = S_f$. In other words, the gravitational force that moves water downstream is balanced with the frictional force acting on the riverbed. With this assumption, the discharge q can be expressed using a uniform flow formula such as Manning's equation:

$$q = A\frac{k}{n}R_\mathrm{h}^{\alpha}S_0^{\frac{1}{2}} \tag{A1}$$

where k is a scalar whose value is 1 for SI units and 1.49 for Imperial units, n is the Manning coefficient, R_h is hydraulic radius [L], which is defined as the cross sectional flow area A [L^2] divided by the wetted perimeter P [L], and α is a constant coefficient ($\alpha = 2/3$).

We assume the channel shape is rectangular and the geometry is constant throughout one river segment, with width w, $A = wy$ and $P = w + 2y$. Assuming the channel is wide compared to flow depth (i.e., $w \gg y$), the hydraulic radius R_h is expressed as

$$R_\mathrm{h} = \frac{A}{P} = \frac{wy}{w+2y} \cong y \tag{A2}$$

By substituting Eq. (A2) into Eq. (A1), the Manning equation is re-written as

$$q = w\frac{k}{n}y^{\alpha+1}S_0^{\frac{1}{2}} \tag{A3}$$

For each stream segment within which the channel width w is constant, the wave celerity C is given by

$$C = \frac{\mathrm{d}q}{\mathrm{d}A} = \frac{\mathrm{d}q}{\mathrm{d}(wy)} \cong \frac{\mathrm{d}q}{w\mathrm{d}y} \tag{A4}$$

By substituting Manning's equation (Eq. A3 into Eq. A4), the wave celerity C can be given by

$$C = (\alpha+1)\cdot\frac{k}{n}\sqrt{S_0}\cdot y^{\alpha} \tag{A5}$$

or expressed as a function of discharge q as

$$C = (\alpha+1)\cdot(w)^{\frac{-\alpha}{\alpha+1}}\cdot\left(\frac{k}{n}\sqrt{S_0}\right)^{\frac{1}{\alpha+1}}\cdot q^{\frac{\alpha}{\alpha+1}} \tag{A6}$$

Appendix B: Derivation of 1-D diffusive equation

We describe in detail the derivation of diffusive wave equations from Saint-Venant equations (Strum, 2001) that are the basis of the IRF-UH method. The development of the IRF-UH method starts with the derivation of the diffusive wave equation from the 1-D Saint Venant equations (Eqs. 3 and 4) by neglecting inertia terms (the second term in Eq. 4) and assuming steady flow (eliminating the first term in Eq. 4). The momentum equation (Eq. 4) can therefore be reduced to:

$$\frac{\partial y}{\partial s} = S_0 - S_\mathrm{f} \tag{B1}$$

Now, Manning's equation can be expressed in terms of channel conveyance, K_c (carrying capacity of river channel),

$$q = K_\mathrm{c} \cdot \sqrt{S_\mathrm{f}} \tag{B2}$$

where $K_\mathrm{c} = k/n \cdot A \cdot R_\mathrm{h}^{\alpha}$. Substituting S_f from Eq. (B2) into Eq. (B1) and differentiating with respect to time, the momentum equation (Eq. A1) becomes

$$\frac{2q}{K_\mathrm{c}^2}\frac{\partial q}{\partial t} - \frac{2q^2}{K_\mathrm{c}^3}\frac{\partial K_\mathrm{c}}{\partial t} = -\frac{\partial^2 y}{\partial x \partial t} \tag{B3}$$

Also, the continuity equation (Eq. 3) can be re-rewritten by differentiating both sides of the equation with respect to distance x as,

$$\frac{\partial^2 q}{\partial x^2} + w\frac{\partial^2 y}{\partial x \partial t} = 0 \tag{B4}$$

Combining Eqs. (B3) and (B4) results in

$$\frac{2q}{K_\mathrm{c}^2}\frac{\partial q}{\partial t} - \frac{2q^2}{K_\mathrm{c}^3}\frac{\partial K_\mathrm{c}}{\partial t} = \frac{1}{w}\frac{\partial^2 q}{\partial x^2} \tag{B5}$$

Because the channel conveyance, K_c, is a function of flow depth, y, or flow area, A, the differentiation part of the second term of Eq. (B5) can be written as

$$\frac{\partial K_\mathrm{c}}{\partial t} = \frac{\mathrm{d}K_\mathrm{c}}{\mathrm{d}A}\frac{\partial A}{\partial t} = -\frac{\mathrm{d}K_\mathrm{c}}{\mathrm{d}A}\frac{\partial q}{\partial x} \tag{B6}$$

Finally, inserting Eq. (B6) into Eq. (B5), results in the one-dimensional diffusive wave equation

$$\frac{\partial q}{\partial t} = D\frac{\partial^2 q}{\partial x^2} - C\frac{\partial q}{\partial x} \tag{B7}$$

where

$$\begin{aligned} C &= \frac{q}{K_\mathrm{c}}\frac{\mathrm{d}K_\mathrm{c}}{\mathrm{d}A} = \frac{\mathrm{d}q}{\mathrm{d}A} \\ D &= \frac{K_\mathrm{c}^2}{2qw} = \frac{q}{2wS_0} \end{aligned} \tag{B8}$$

where parameters C and D are wave celerity $[\mathrm{L\,T^{-1}}]$ and diffusivity $[\mathrm{L^2\,T^{-1}}]$, respectively. Here, we assume the flow is uniform (i.e., $S_\mathrm{f} = S_0$).

Code availability

The source codes for the river network topology program and the hillslope and river routing along with test data are available along with the user manual on GitHub (https://github.com/NCAR/mizuRoute). Those codes are developed in Fortran90 and require installation of a NetCDF 4 library (http://www.unidata.ucar.edu/downloads/netcdf/index.jsp). In addition, there are several pre-processing python scripts to map runoff outputs from hydrologic models to other type of HRUs. These pre-processing scripts are also available in GitHub. Those python scripts process ESRI Shapefiles and NetCDF data and require GDAL, SHAPELY, NetCDF4 package.

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Title Page
Abstract Introduction
Conclusions References
Tables Figures
◀◀ ▶▶
◀ ▶
Back Close
Full Screen / Esc
Printer-friendly Version
Interactive Discussion

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Table 1. Routing model parameters.

Parameters	Routing methods	Descriptions	Values used in Sect. 5
a	Hillslope	Shape factor [–]	2.5 [–]
θ	Hillslope	Time scale factor [T]	86 400 [s]
n	KWT	Manning coefficient [–]	0.01 [–]
W	KWT	River width scale factor [–]	0.001 [–]
C	IRF-UH	Wave velocity [$\mathrm{L\,T^{-1}}$]	1.5 [$\mathrm{m\,s^{-1}}$]
D	IRF-UH	Diffusivity [$\mathrm{L^2\,T^{-1}}$]	800 [$\mathrm{m^2\,s^{-1}}$]

Table 2. River network information required in mizuRoute.

Variables	Vector data type	Descriptions
seg_id	River segment line	ID of segment
tosegment	River segment line	ID of immediate downstream segment
Length	River segment line	length of segment [m]
Slope	River segment line	Slope of segment [$\mathrm{m\,m^{-1}}$]
hru_id	HRU polygon	ID of HRU
Seg_hru_segment	HRU polygon	ID of segment to which the HRU discharge
hru_area	HRU polygon	Area of HRU [$\mathrm{m^2}$]

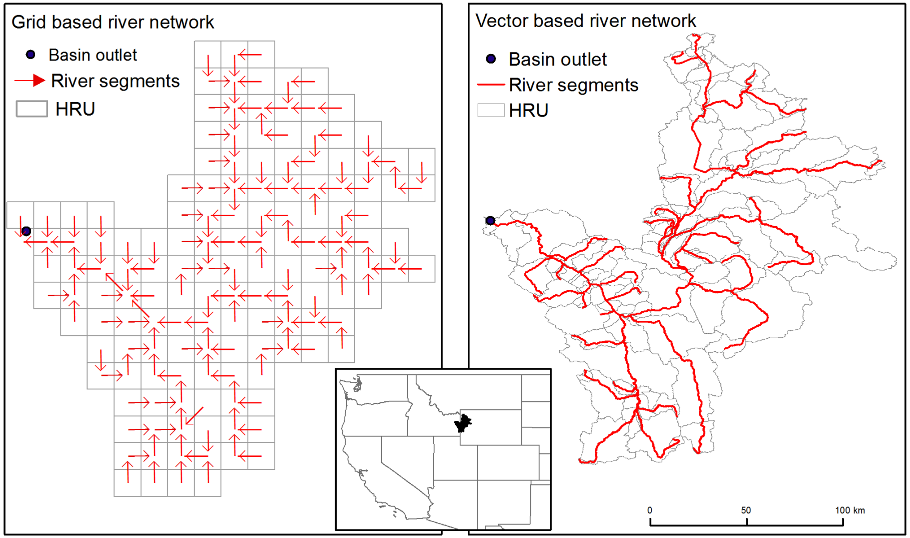


Figure 1. Comparison of 1/8th degree (∼ 12 km) gridded river network and vector river network from United States Geological Survey (USGS) Geospatial Fabric for the upper part of Snake River basin (Vigor, 2014).

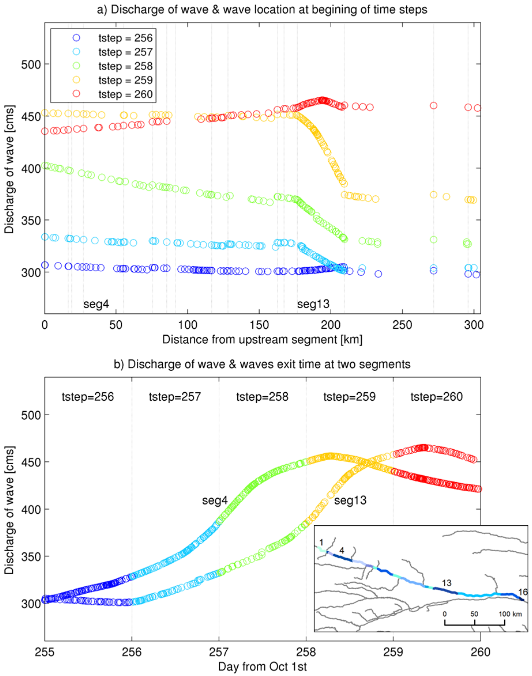


Figure 2. Visualization of waves. The top panel **(a)** plots a discharge of the wave (cms) against its location (distance (km) from the beginning of the 1st segment) at the beginning of five consecutive time steps. A vertical line indicates the river segment boundary. The bottom panel **(b)** plots a discharge of the wave (cms) against its exit time (day from 1 October) for the 4th and 13th segments. A vertical line indicates the boundary of the routing time step. The inserted map shows the 16 river segments used for the plots.

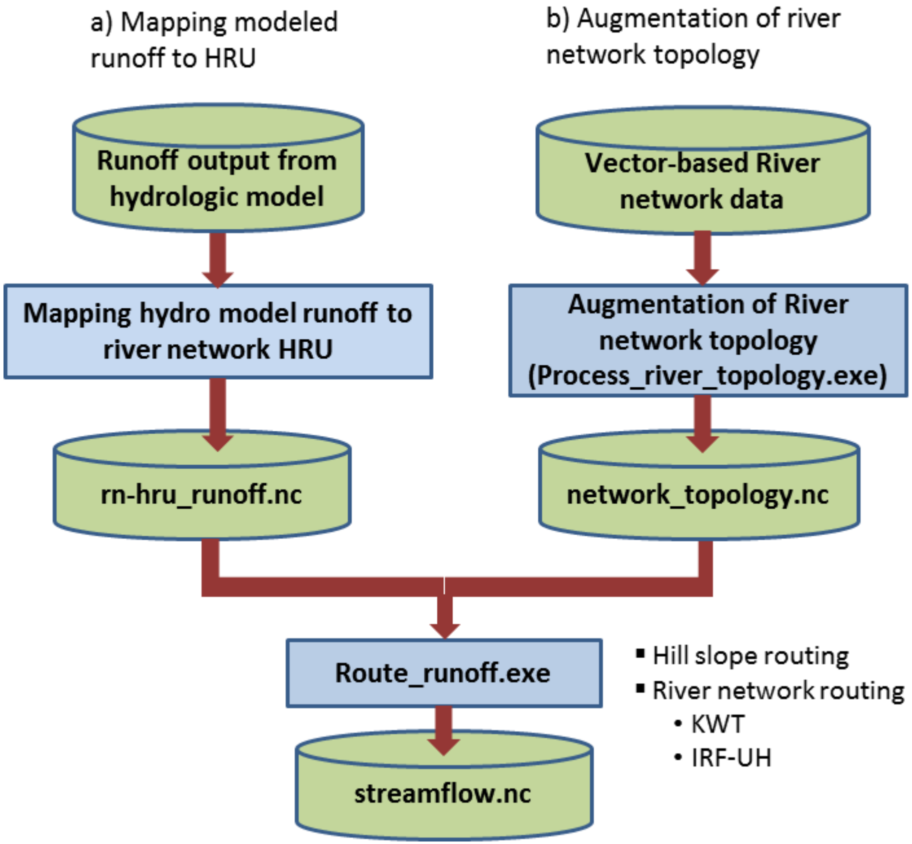


Figure 3. Overview of streamflow simulation with mizuRoute.

Title Page
Abstract Introduction
Conclusions References
Tables Figures
◀◀ ▶▶
◀ ▶
Back Close
Full Screen / Esc
Printer-friendly Version
Interactive Discussion

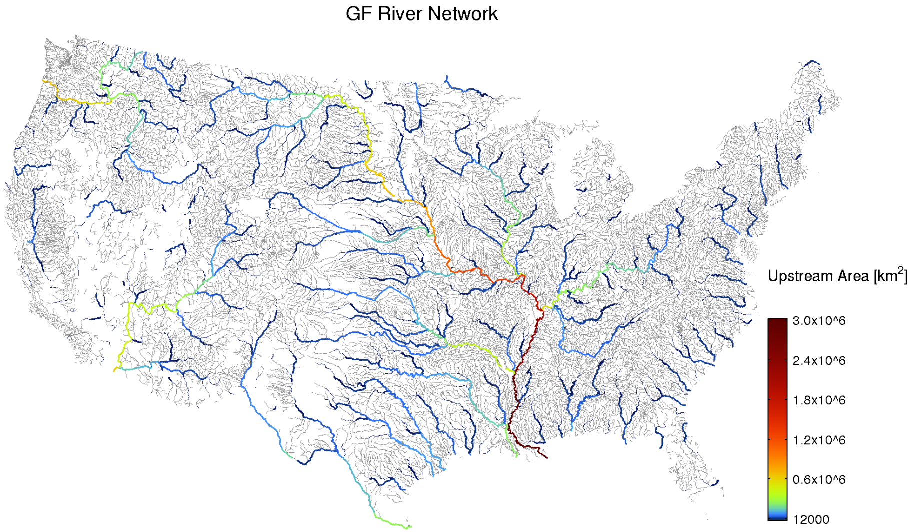
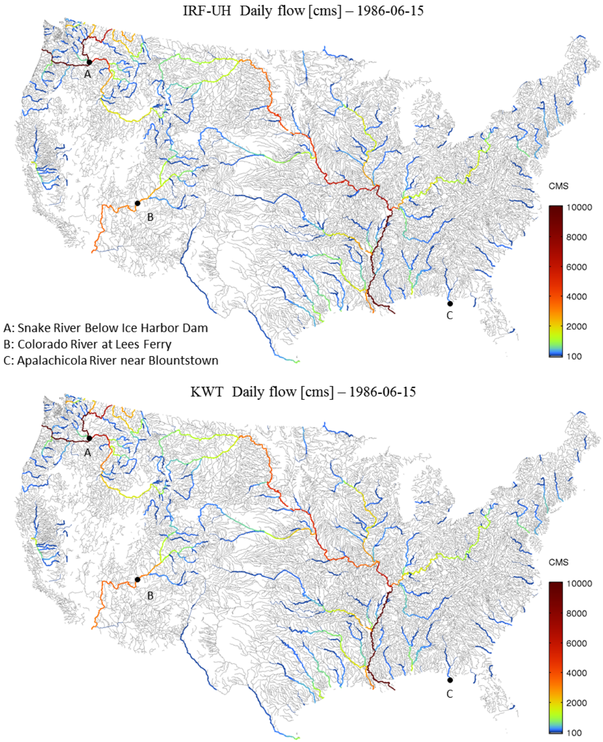


Figure 4. GF river network color coded by upstream drainage areas. Gray lines indicate the total upstream drainage areas less than 12 000 km^2.

Title Page
Abstract | Introduction
Conclusions | References
Tables | Figures
◀| |▶
◀ ▶
Back | Close
Full Screen / Esc
Printer-friendly Version
Interactive Discussion

Title Page
Abstract
Introduction
Conclusions
References
Tables
Figures
◀▌ ▶▌
◀ ▶
Back
Close
Full Screen / Esc
Printer-friendly Version
Interactive Discussion



IRF-UH Daily flow [cms] – 1986-06-15

KWT Daily flow [cms] – 1986-06-15

Figure 5. Spatially distributed daily streamflow on 15 July 1986 in the GF river network simulated with mizuRoute. Gray lines indicate flow less than 100 cms. The streamflow time series are extracted at three USGS gauges (A–C) shown in Fig. 5.

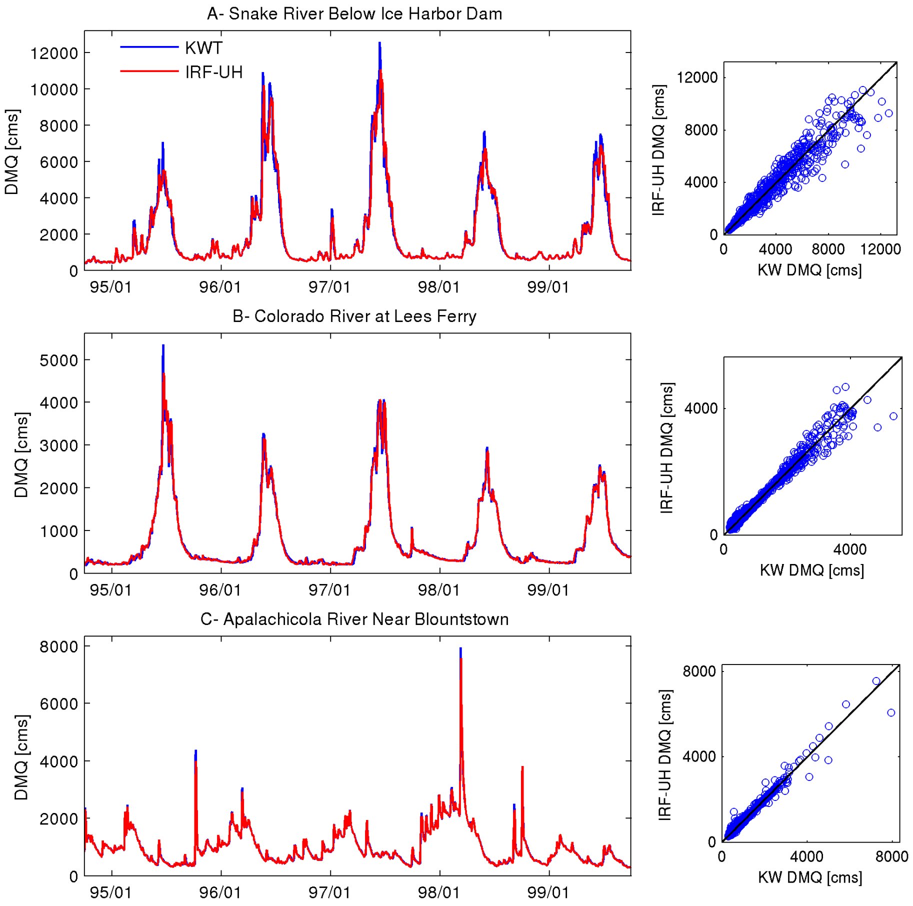


Figure 6. Daily mean streamflow (DMQ) at selected location in the GF river network. A–Snake River below Ice Harbor Dam, B–Colorado River at Lees Ferry, and C–Apalachicola River near Blountstown from top to bottom. The locations of the three gauges are shown in Fig. 4.

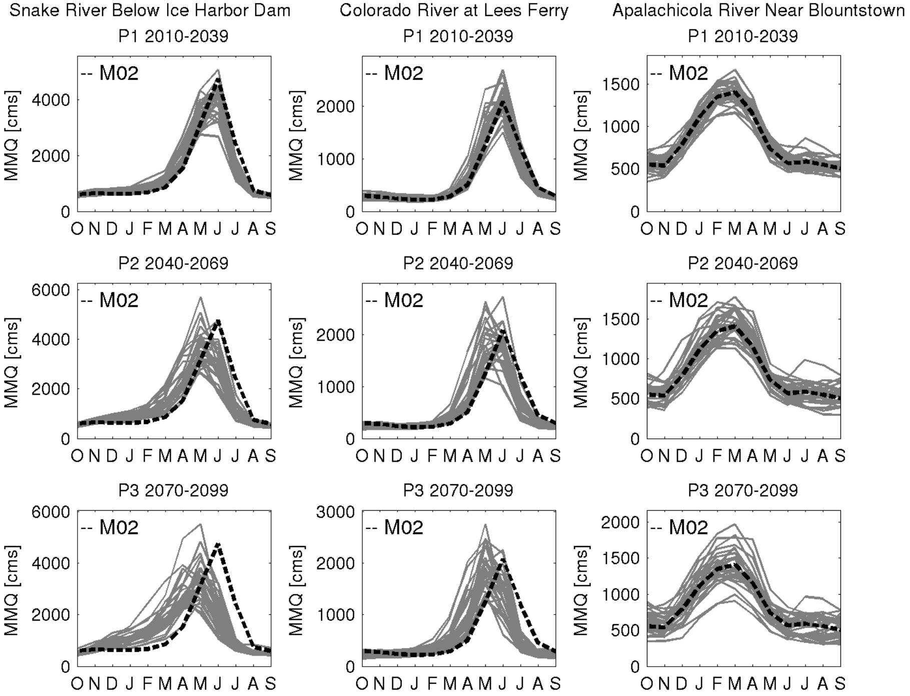


Figure 7. Monthly mean of CMIP5 projected streamflow at three locations indicated in Fig. 4. (Left column – Snake River Below Ice Harbor Dam, middle column – Colorado River at Lees Ferry, and right column – Apalachicola River Near Blountstown). The river routing scheme is KWT. Monthly mean values are computed over three future periods (Top – P1 2010-2039, Middle – P2 2040–2069 and Bottom – P3 2070-2099). The dashed line denotes streamflow estimated from runoff output from VIC forced by M02 historical data while grey lines indicate projected streamflow based on future runoff outputs from VIC forced by 28 CMIP5 RCP8.5 data.

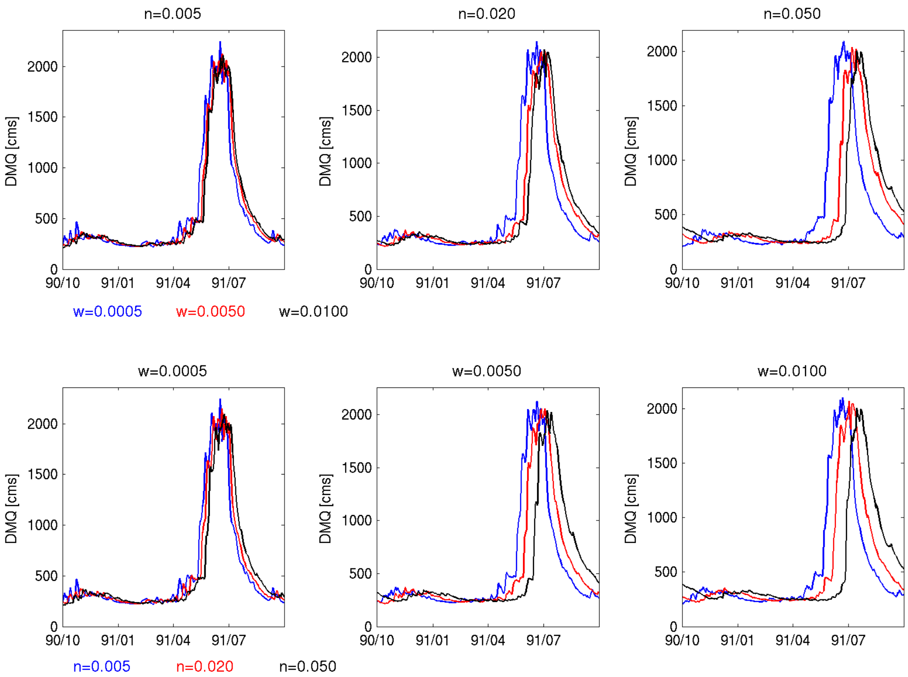


Figure 8. Sensitivity of simulated runoff at Colorado River at Lees Ferry (Location B in Fig. 4) to the two KWT parameters. The top panel shows sensitivity to width factor *w* with three fixed Manning coefficients *n* (from left to right: $n = 0.005$, 0.02, and 0.05). The bottom panel shows sensitivity to Manning coefficient *n* with three fixed width factor *w* (from left to right: $W = 0.0005$, 0.0050, and 0.0100).

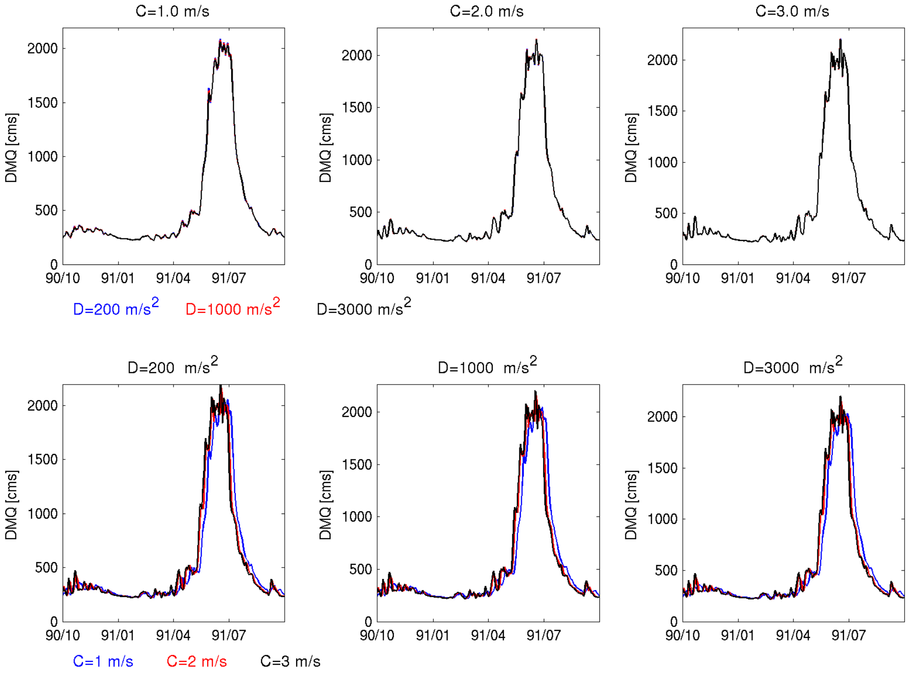


Figure 9. Sensitivity of simulated runoff at Colorado River at Lees Ferry (Location B in Fig. 4) to IRF-UH parameters. The top panels show sensitivity to diffusivity D with three fixed celerity C values (from left to right: $C = 1.0$, 2.0, and 3.0 $\mathrm{m\,s^{-1}}$). The bottom panels show sensitivity to celerity C with three fixed diffusivity D values (from left to right: $D = 200$, 1000, and 3000 $\mathrm{m\,s^{-2}}$).