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Dear Jeremy,

Please find attached our revised paper with changes from the original version tracked. We found both the reviewers' and the interactive discussion comments on the manuscript to be very helpful and have responded in detail below. In particular, we have revised the design of the core experiment to include freshwater forcing from ice sheets in line with the reviewers' and Eric Wolff's comments. We have also emphasised the scientific purpose of the experiment; to understand the sequence of last deglaciation events using 'model-geological data' comparison.

We wish to highlight that the Core experiment has been designed following extensive consultation (since January 2014) across the palaeoclimate modelling and reconstruction community. It has been challenging to reconcile everyone's preferences for such a demanding experiment (more than 12 thousand model years); everyone uses different tools and has different (though complementary) scientific aims to each other. However, we are confident that in its new form, the Core experiment design represents the best compromise between any conflicting views and leaves room to accommodate everyone's priorities and resource limitations. For example, this is one reason why we have to be flexible on the ice sheet meltwater forcing protocol, including allowing groups to run without meltwater if they wish (although we recommend that they run with meltwater and provide scenarios that are consistent with the ice sheet histories). Certainly these efforts will be worthwhile and we are excited to begin the experiment and see the first results.

Reviewer 1

Reviewer's summary:

The paper describes the design of the coordinated Core simulation over 21-9 ka with time varying orbital forcing, greenhouse gases, ice sheets and other geographical changes. The choice of two ice sheet reconstructions is given but no meltwater is prescribed. The paper reviews in detail the past experimental designs by EMICs and

AOGCMs (ex CCSM) and their results but unfortunately mismatches the experimental design presented this time because of no meltwater. I am afraid the readers are lost in understanding what we can learn from the experiment at the present form of the paper. I recommend the paper published after revision by (1) [equivalent to point 1. below] presenting a core experimental series with melt-water given at least in a very simple way and (2) [equivalent to point 2. below] explaining what kind of analysis are useful after collecting the non-meltwater experiments from PMIP community. Also (3) [equivalent to point 3. below] clarify the design related to coastline, bathymetry and salinity change due to ice sheet change.

- 1. Reviewer's comment:** '(1) For the first point, the meltwater that is consistent with the ice sheet provided from two schools should be provided so that additional experiment with meltwater can be performed. Where to release could be an option. As in many studies, the regional difference (South vs. North) during the deglaciation is presented and discussed but without the meltwater there is no way expecting the reproduction in proxy as in Clark et al, 2012 or Shakun et al, 2012. Even if there is uncertainty of the location of the meltwater or an uncertainty of timing of abrupt change of melt water, at least the total amount of meltwater can be provided and given by each modelling group. The change of total amount (~ sea level change) should be consistent with the ice sheets reconstructed and also constrained fairly well (Clark et al, 2009).'

Authors' response: We have updated the Core experiment design to include freshwater and have revised the manuscript accordingly. This includes uniformly distributed freshwater to conserve water in the simulation (i.e. the total amount of ice melt is applied uniformly to the ocean) and 'routed' freshwater (i.e. fluxes from particular coastal outlets, which can be used to examine more regional responses); both are consistent with the ice sheet reconstructions provided (GLAC-1D and ICE-6G_C). Further *focussed* simulations will explore these hypotheses (e.g. north vs. south meltwater injection sites) more thoroughly and systematically.

- 2. Reviewer's comment:** '(2) For the second point, if the PMIP4 Core experiment group asks for the non meltwater experiment, then the reason and what is expected should be described clearly. If there is no melt-water, there is no sense in doing a transient experiment, which is very expensive. It is unclear at the moment why the non-melt water experiment should be done as a Core experiment. PMIP experiments with AOGCMs are expected to do model-data comparison [sic] as well as model-model experiment, but what are the data-model comparison expected? Many studies suggest that the melt water might be important for understanding the "bipolar" ice core signals and various regional signals in proxy. Since the experiment demands substantial computational resource as well as man-power for many groups, the explanation should be convincing. The introduction in the paper is not sufficiently written for the non-melt water transient experiment. Perhaps what is expected scientifically after collecting the results could be written in an independent section in more detail.'

Authors' response: As addressed above (point 1), we have taken on board this comment and revised the manuscript and experiment design to recommend including meltwater (consistent with the ice sheet reconstructions) in the Core. We have also emphasised the aim to carry out model-data comparison with the results (e.g. section 1.4), for which including meltwater forcing in the Core simulation is preferable. We have expanded section 2.5 to discuss this in more detail, including the value of running meltwater-free simulations to accompany meltwater-include simulations.

3. **Reviewer's comment:** '(3) On the design related to coastline and bathymetry change due to ice sheet change: In table 2, the design of salinity change is unknown. Define what (and how) the modelers do with the total ocean salinity change, which should be consistent with the ice sheet change and melt water.'

Authors' response: We have expanded the row on Freshwater fluxes' in table two to include the addition of freshwater to the Core and to explicitly advise groups to conserve salinity. The freshwater scenarios provided will conserve salinity changes relating to ice sheet evolution (following GLAC-1D and ICE-6G_C). We cannot be more precise about how this will be implemented (technically) because it is model-specific and therefore up to the user, but the transient data will be provided to make this possible.

Reviewer's comment continued: 'There are two options for the "Bathymetry" but what happens when the ice sheet covers the ocean in the model that keeps the "Preindustrial bathymetry"? What is prescribed for ice sheet and what should be done for ocean boundary condition should be carefully designed and described for the participants.'

Authors' response: The land-sea mask (or 'coastlines') will need to be consistent with the ice sheet, as outlined in the text and tables. This means that land should underlie the grounded ice sheets, as requested by the reviewer, but does not require other bathymetric changes. We have extended the text and table entry to clarify this in the manuscript. Again, how this is technically implemented is model-specific and best decided by the expert user.

4. **Reviewer's comment:** 'Page 9073 line 9 "many questions and untested hypotheses remain" but the current study should show the perspective, how it answers the questions and the hypotheses are tested.'

Authors' response: A large component of section 1 is devoted to showing the context (or perspective) of the working group, outlining the current state of our knowledge, what hypotheses exist and what questions remain. We have extended the text in this section, as well as later sections, to explain how the multi-model approach aims to narrow down uncertainty in the ice sheet reconstruction and meltwater forcing, for example. We have also discussed *focussed* simulations that will represent rigorous sensitivity- and hypothesis-driven investigations that are of particular interest to participant groups; including, for example, the regional specificity of climate system response to freshwater

inputs, the timing of changes in greenhouse gas records, the influence of the acceleration of northern ice melt on ocean circulation during Heinrich Stadial 1 – these are all discussed in the manuscript and we have extended the text in relevant sections to clarify this. These foci will be further defined on the last deglaciation PMIP Wiki and in subsequent manuscripts as they are investigated.

Reviewer 2: Shawn Marshall

Reviewer's summary: 'Summary This manuscript describes the scientific motivation and technical specifications for a community model experiment simulating the deglaciation (26 or 21 ka until 9 ka BP) in climate models of differing complexities. The experiments are designed so that both fully- coupled Earth system models and a variety of reduced models can take part. There is a nice blend of flexibility in the model design – with specified boundary conditions for the main climate forcings and their temporal variability, but some user discretion on implementation. The balance seems appropriate. This is nicely presented and explained, overall. The summary of deglacial climate dynamics and some of the paleoclimatic enigmas during this period makes for a lovely review, and the experiments that are described will certainly be interesting. Most of what one needs from this manuscript is encapsulated in Tables 1 and 2, so at first it seemed unnecessarily long, but the narrative is nicely written and a pleasure to read, offering some helpful insights about the approach to be adopted in the intercomparison.'

5. Reviewer's comment: 'I am a bit surprised that the 'focussed' experiments are not described or prescribed in detail at this point. I understand that perhaps these need to be reactionary to the results of the core experiment. It seems unfortunate though, as it would be helpful to have this information together in a single document. I am sure lots of ideas are already in place for the spinoff or focussed experiments, and it would not have taken too much extra work to have these set out here. But this is not necessary, and it is probably helpful to keep these flexible and as subsets of the main modelling exercise.'

Authors' response: The Core itself is a significant undertaking for modelling groups, it has taken a lot of discussion to agree upon the experiment design within the community. Also, different groups have different preferences and priorities for the *focussed* simulations. Therefore we agree with the reviewer that it is helpful at this stage (and until the Core is published) to keep the *focussed* experiment designs flexible, and for clarity it is preferable to keep their full specification separate from the required Core (described in detail here). As does the reviewer, we also expect some of the *focussed* simulations to be reactionary to working group results, as they emerge.

However because some discussions are already underway and, as the reviewer suggests, we do have some ideas in place, we have extended the text in sections 2.5 and 3 to provide more information on currently planned *focussed* simulations.

6. Reviewer's comment: 'My only substantive feedback or suggestion involves the meltwater treatment. Several thoughts related to this are made below, in the specific comments. Overall, it seems inconsistent to have specified, time-varying ice sheet volume on the continents but not honour this global water conservation when it comes to the ocean freshwater and salinity budget. I appreciate the desire to control for meltwater runoff, but it makes one wonder if the core experiment, as described, is meaningful since it does not do a physically sensible job of representing the basic ocean state through deglaciation. At least as I understand the model design. Things like preconditioning and ocean mixing surely depend on the mean salinity and its structure. I appreciate that this design is intentional, to eliminate some of the complexity and model dispersion associated with when/where to put the meltwater. And models are dealing with meltwater routing and runoff internally, in some cases. But since the specification is to violate water balance and neglect runoff processes, it would not be unreasonable to honour water balance while neglecting runoff processes. That is, the ice sheet \dot{V}_{ADV} , [sic] as specified through the Peltier or Tarasov reconstructions, can be converted to eustatic water equivalent and restored to the nearest ocean in a specified way for all model experiments. This could be considered for the Core experiment as something a bit more realistic, while saving some of the detailed questions about meltwater runoff and iceberg discharge for the focussed experiments. It would require a bit of extra work to define the timing and location of freshwater runoff, which everyone would follow, but this can be straightforward I think. Just don't inject the water all at once every 1000 years, when the ice geometry changes. Rather than shock the system, one could, for example, take the 1000-year \dot{V}_{ADV} in each major river catchment and divide \dot{V}_{ADV} by 1000 to give the average runoff in m^3/yr (or convert to Sv), in a way that respects water balance. If one wants to avoid some of the detailed questions concerning paleoriver routing, the appropriate amount of water could just be spread over the large-scale basin (e.g. North Atlantic, Southern Ocean, etc.). I would leave it to the authors to consider what is best here, but I do recommend considering a treatment like this within the core experiment design.'

Reviewer's comment: 'p.9047, ll.12-14, "A choice of two ice sheet reconstructions is given, but no ice sheet or iceberg meltwater should be prescribed in the Core simulation." – this is confusing, are ice sheets to be prescribed or internally modelled? I understood what the authors meant by the end of the manuscript, i.e. don't put any ice sheet meltwater into the oceans, but this seems contradictory to prescribe ice sheets but not put the prescribed change in water volume back into the oceans.'

Reviewer's comment: 'p.9067, Section 2.5, freshwater fluxes during the deglaciation. It does seem odd but also sensible to have controlled experiments that examine non-meltwater forced climate change during the deglaciation. Although given the important role that ocean circulation simply had to have played in the Bolling and YD, this seems limiting. i.e., orbital forcing and CO₂

clearly cannot explain these features of the deglaciation. A reference experiment is nonetheless important and useful. I wonder if it is the best reference though, given that the ice sheets did melt away and ocean salinity did decrease through this period. Is it possible to have prescribed changes in mean ocean salinity through the deglaciation and/or prescribed runoff as a second core experiment? The latter could be done based on the 1000-yr ice sheet updates to at least have the correct global water cycle (conservation). I appreciate the arguments and intricacies concerning when and where to put the meltwater. Some hypothesis-driven experiments here seem sensible, as additional experiments.'

Authors' response: We have taken on board this feedback and have adapted the Core experiment design to include transient (i.e. not stepped/shocked) meltwater fluxes in the Core simulation (also see reply to reviewer comment 1 and 2, and the second paragraph of this letter). This will enable more fruitful model-data comparison and the possibility to narrow down uncertainty in last deglaciation ice sheet meltwater fluxes. Our *focussed* simulations (e.g. as briefly outlined in section 3) will continue to address this more fully.

7. **Reviewer's comment:** 'p.9048, l. 26, "majority of its ice melting" – not really the majority of the Antarctic Ice Sheet melting; rather, much of the excess LGM ice that was out on the shelf, and the thicker ice that covered WAIS; but overall, it was closer to a 20% loss of the ice in Antarctica through this period'

Authors' response: We've changed this sentence in line with this comment to make the meaning clearer.

8. **Reviewer's comment:** 'p.9050, l.4, the idea of mid-latitude N.Atlantic warming during H1. This is not really compatible with the preservation of Hudson Strait icebergs in a swath at 40-55 N across to Portugal. Is it more of a subtropical warming that has been proposed? Else it is perhaps worthwhile to note this incompatibility.'

Authors' response: We have removed this last clause in the sentence.

9. **Reviewer's comment:** 'p.9050, l.15, suggest deleting 'older', it conveys a bias against these studies, i.e. a potential lack of objectivity, whereas many of the studies cited below in favour of a northern source are in fact older'

Authors' response: We have removed 'older'.

10. **Reviewer's comment:** 'p.9062, Section 2.1. I wonder about a prescription for oceanic or surface ocean d18O and dD as well, for those that will explore isotopic cycles through the deglaciation.'

Authors' response: This is beyond the scope of the Core simulation, which is designed as a basic simulation for all models (and most will not run with water isotopes, $\delta^{18}\text{O}$ and deuterium, due to the computational expense or because they are not implemented), but it would make an appropriate *focussed* experiment. We will discuss this in the group and if it is popular, we will work in close collaboration with the *Isotopes* PMIP working group to design the experiment, including the prescription of water isotopes. It will require careful planning because the isotopes are implemented in the different models in different ways, so we would need to provide the most valuable and important data. For example, many models will get their ocean surface water isotopes through interaction with the atmosphere as the isotopes are implemented throughout the whole hydrological cycle. However, changes in terrestrial ice volume that are not dynamically simulated in the model will complicate this, and that may be where we will need to provide a transient global budget of water isotope data and meltwater signatures.

- 11. Reviewer's comment:** 'p.9062, II.7-9, discussion of the freshwater budget. Just to be clear here, the experiments should prescribe/force all precipitation to return to the oceans annually then, i.e. equilibrium mass balance conditions on the ice sheets? This is fair for present purposes, but I guess that it will not occur naturally in any of the models, so this sounds tricky. I wonder if more explicit directions here would be helpful, as to how the freshwater routing/flux adjustments should be prescribed. For instance, should an LGM catchment map be prescribed, so that everyone is using the same one, based on the ice sheet configuration? Then everyone forces all precipitation within the catchment to return via a prescribed river outlet/coastal grid cell.'

Authors' response: This is a technical point specifically relating to the equilibrium-type spinup of the LGM to make sure that during the spinup, there are no large salinity drifts in the model, and that water is conserved. It is difficult to provide more detailed or precise directions because it is so model-specific. However, it has been common PMIP practice for several years (e.g. for the LGM experiment) to have to consider this (in the LGM spinup); we have provided the most recent text from the PMIP Wiki, but similar earlier advice was given (e.g. PMIP3 and CMIP5). This should be sufficient information, but if not, individuals can use the working group mailing list and Wiki to ask for help in generating their spinup; although support from their model developers is likely to be more useful in this instance.

Similarly, LGM catchment maps will be provided as they become available. However, when these have been offered in the past, uptake has been very low due to the technical challenges involved (we are not aware of any groups having implemented it). Furthermore, some models calculate their own river routing based on surface topography. For these reasons, we will provide the data, but leave the choice of what to implement for river routing to the expert model-user. However, it is essential to ensure that rivers reach the coast, and this is explained in the manuscript (section 2.6). These details are provided by the *LGM* PMIP working group.

12. Reviewer's comment: 'p.9064, ll.23, 27. I think with Tarasov as an author, you don't have to list this as 'personal communication' – also on the next page'

Authors' response: We have amended these lines.

Interactive Discussion

13. Comment by A. Carlson: 'Hi all, so just looking through the text, I noticed one incorrect statement on page 9055 lines 18-22. The timing of CIS-LIS separation was around the time of MWP-1A according to Dyke (2004) is not right. Dyke (2004) specifically states: "Unfortunately, the initial opening of the ice-free corridor remains only tenuously dated. The initial opening at the south end probably started about 15 ka BP, based on an AMS date of 15.67 ka on wood, mentioned above (Beierle & Smith, 1998) and exposure dates on the Laurentide terminal moraine in the south-western Alberta Foothills (Jackson et al., 1999). By 13.5 ka BP, the southern half of the corridor seems to have opened, because wood (evidently small wood, probably arctic willow; P. Bobrowsky, personal communication, 2002) from basal sediment of glacial Lake Peace yielded an AMS date of 13.97 ka BP (Catto et al., 1996). It is possible that the corridor was entirely open by 13.5 ka BP, because the dated site is located midway within it. However, in the absence of similarly old dates from the northern part of the corridor, and assuming slower ice ablation further north, initial joining of the southern and northern approaches of the corridor may not have occurred until 12.5 ka BP or possibly even 12 ka BP. It seems exceedingly unlikely, however, that ice coalescence could have continued until 11.5 ka BP, for the Mackenzie Lobe of Laurentide ice had by that time receded halfway up the Mackenzie Valley (Mackay & Mathews, 1973; Smith, 1992). Furthermore, south-eastern Cordilleran ice had by then receded into the alpine zone (Reasoner et al., 1994), and Cordilleran ice distribution in Alaska was close to the present one. In summary, the known history of the ice-free corridor, although imprecise, does not preclude the possibility of pre-Clovis people using this route and its availability to early Clovis people is almost certain." These are all ages in 14C years, meaning the corridor started opening by ~18.8 cal ka and was over half complete by ~17 cal ka, thousands of years before MWP-1A. I think this section needs to be clearly redone to reflect this much slower and earlier separation of the CIS-LIS that Dyke (2004) discussed to stop'

Authors' response: As is pointed out in this comment, an adequate discussion of the issues raised would be too lengthy and is only peripheral to this manuscript. We have thus removed this sentence from the manuscript.

14. Summary by EW Wolff: 'This is not mainly a review of the paper but is, as requested by the lead author, a set of comments on the proposed experiment design. In general of course this is a usefully comprehensive description of what is planned under PMIP for the deglaciation transient. I have a few rather minor comments and then one that is more significant.'

Page 9049, line 11. This paper should be referred to as EPICA Community Members, 2004 rather than Augustin et al 2004.'

Authors' response: We have made this change.

15. Comment by EW Wolff: 'Page 9050, line 9. It is a little misleading to say that a shift in climate occurred in 1-3 years. A rapid shift occurred in some components (d-xs most notably) but for example the inferred temperature change was slower. I suggest just adding "some components of" climate.'

Authors' response: We have made this change.

16. Comment by EW Wolff: 'Page 9057. Should you add that an important challenge for PMIP is to assemble suitable datasets for model-data comparison. Probably you say that elsewhere.'

Authors' response: Yes, we've added text to this effect at the end of the first paragraph of section 1.3

17. Comment by EW Wolff: 'Page 9062, line 8 [sic – page 9063 line 12?] and numerous other places, including Table 1. I am sure you mean "i.e.", meaning "that is", and not "e.g." meaning "for example". This is important as I assume you are telling participants they must use 1365 W/m², not that they can use any number they consider represents the preindustrial?'

Authors' response: In the case of the solar constant, it is deliberate to use 'e.g.' (for example) because the recommended preindustrial value (1365 W m⁻²) is a widely used and accepted value, it has thus far been the recommended PMIP value, but is not exclusively accepted and others may also be used (see recommendations at <http://solarisheppa.geomar.de/cmip5>). Some groups already prefer to use other PI values; e.g. 1370 W/m² (Zhang et al., 2012) and 1360.9 W/m² (Landrum et al., 2013).

18. Comment by EW Wolff: 'Page 9063, line 4, and other places including Tables 1 and 2. You suggest using the Luthi et al 2008 data (which for this part of the core is really the Monnin et al 2001 data) translated to AICC2012. This is an option, but you might want to at least discuss using the dataset presented in Bereiter et

al (2015) as supplementary data. Here they have already done the work of translating to AICC2012, and they include a range of datasets in their composite dataset, including the high resolution WAIS Divide data, with a 4 ppm offset (the offset discussed later on page 9063). To me it would seem smarter to use the fully resolved but consistent dataset.'

Authors' response: We have updated the experiment design, manuscript text figures and references to use the more recent Bereiter et al. (2015) CO₂ data.

19.Comment by EW Wolff: 'Page 9067, last paragraph. "Can abrupt deglacial changes be simulated without icemeltwater?". I think this is a bit disingenuous. We already know that they can't: the north-south phasing of climate is simply wrong if freshwater is excluded as already shown clearly in papers including Shakun et al (2012).'

Authors' response: We have added freshwater to the core simulation.

20.Comment by EW Wolff: 'Page 9069. Regarding dust, isn't this another parameter that might be varied in extended simulations?'

Authors' response: Yes, we have extended the text to include this suggestion.

21.Comment by EW Wolff: 'Page 9071, line 13. Do you mean "timing" in comparing Luthi to Marcott. I think we can easily fix any timing mismatches, as done in Bereiter et al (2015); it is really resolution that is the issue.'

Authors' response: We have removed the comparison of Luthi et al. (2008) to Marcott et al. (2014) since this can be resolved in the records.

22.Comment by EW Wolff: 'Page 9072. "the ...design for later periods..is updated". I don't really see how this will work. Some groups will quite sensibly run straight through the whole period. It will be very confusing if you then change some aspect of the design halfway through, just because others have now reached a milestone. Are you really suggesting groups should hold their simulation at the end of each phase until everyone reaches the same point?'

Authors' response: This would be important to avoid (as is pointed out, 'some groups will quite sensibly run straight through the whole period') so we have clarified

this point in the text (first paragraph of section 4). Mainly, changes will not compromise the Core, although new data may be used to design *focussed* spin-off simulations, and to assess the Core results in light of the changes (and additional simulations).

23. Comment by EW Wolff: ‘My major comment comes back to what the purpose of the experiments is, as always with PMIP. I can see two main classes of justification. One is to test different models against data. The other is to compare the performance of different models against each other. If the aim is the former then it makes sense to allow people some freedom to use different boundary conditions, which you do in allowing two different ice models. If it’s the latter it makes no sense to have radically different ice models. However it cannot be the former, because you already know that in the core experiment, you won’t get anything like the data (because no bipolar seesaw contrast). Given that, the core experiment (but not the extended ones) **MUST** be aiming mainly at model-model comparisons and these can only be made if most features of the design are common. I realise you probably had groups who would not compromise on use of their favourite ice model, and I sympathise with the dilemma but not the solution. I think you have to be firm and choose a primary ice model, with no suggestion that it is better and with a strong recommendation that as many groups as possible run both. Those who want to use whichever you choose as the secondary ice model can use it as long as they also use the primary one in a parallel experiment. The aim should be to have a situation where the model-model comparison an [sic] be made without compromise.’

Authors’ response: We have carefully considered this point and have (a) adapted the Core to include meltwater, and (b) emphasised the importance of model-data comparison in the working group’s aims (some of the *focussed* sensitivity experiments will enable model-model comparisons, but our main priority is model-data comparisons). Also see response to Reviewer comment 1, 2 and 6 above.

24. Comment by A. Schmittner: ‘In a recent paper (Schmittner et al. 2015) we have shown that changes in tidal energy dissipation between the LGM and the late Holocene may have a large impact on the Atlantic Meridional Overturning Circulation. I think it may be warranted to think about if this could be included as a prescribed forcing over the deglaciation. Schmittner, A., Green, J. A. M., and Wilmes, S.-B. (2015) Glacial Ocean Overturning Intensified by Tidal Mixing in a Global Circulation Model *Geophysical Research Letters*, 42 (10), 4014-4022. doi: 10.1002/2015GL063561
<http://onlinelibrary.wiley.com/doi/10.1002/2015GL063561/full>’

Authors' response: This will be difficult to include in the Core design, but is an interesting component to consider. We will propose it as a theme for a *focussed* experiment and have amended the text in section 3 accordingly.

Having carefully and thoroughly addressed all of the reviewers' comments and the interactive discussion, we hope that the revised paper is now acceptable to be published in *GMD*.

Yours sincerely,



Ruža F. Ivanović

References cited in response

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Transient climate simulations of the deglaciation 21-9 thousand years before present; PMIP4 Core experiment design and boundary conditions.

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Abstract

The last deglaciation, which marked the transition between the last glacial and present interglacial periods, was punctuated by a series of rapid (centennial and decadal) climate changes. Numerical climate models are useful for investigating mechanisms that underpin the [climate change](#) events, especially now that some of the complex models can be run for multiple

millennia. We have set up a Paleoclimate Modelling Intercomparison Project (PMIP) working group to coordinate efforts to run transient simulations of the last deglaciation, and to facilitate the dissemination of expertise between modellers and those engaged with reconstructing the climate of the last 21 thousand years. Here, we present the design of a coordinated Core ~~simulation~~experiment over the period 21-9 thousand years before present (ka) with time varying orbital forcing, greenhouse gases, ice sheets, and other geographical changes. A choice of two ice sheet reconstructions is given, ~~but not~~and we make recommendations for prescribing ice ~~sheet or iceberg~~ meltwater ~~should be prescribed(or not)~~ in the Core ~~simulation~~experiment. Additional *focussed* simulations will also be coordinated on an ad-hoc basis by the working group, for example to investigate more thoroughly the effect of ice ~~sheet and iceberg~~ meltwater, ~~and on climate system evolution, and to examine~~ the uncertainty in other forcings. Some of these *focussed* simulations will ~~focus on~~target shorter durations around specific events ~~to in~~ order to understand them in more detail and allow the more computationally expensive models to take part.

1 Introduction

1.1 Climate evolution over the last deglaciation

The last deglaciation is a period of major climate change, when Earth transitioned from its last full glacial state, to the current interglacial climate. The *Last Glacial Maximum* (LGM) marked the culmination of the last glacial cycle when vast ice sheets covered large regions of the Northern Hemisphere, stretching over North America and Eurasia (e.g. Boulton et al., 2001; Dyke et al., 2002; Peltier et al., 2015; Svendsen et al., 2004; Tarasov et al., 2012), and the Antarctic Ice Sheet expanded to the edge of the continental shelf (Argus et al., 2014; Briggs et al., 2014; Lambeck et al., 2014 and references therein). Changes in the ice sheets resulted in a total sea level rise of ~115-130 m between LGM and the late Holocene (Lambeck et al., 2014; Peltier and Fairbanks, 2006) depending upon the time assumed to correspond to the LGM-, and ~100 m from 21 ka to 9 ka (the period of focus for this manuscript).

Historically, the EPILOG group defined the LGM as having occurred 23-19 ka (21 ka centre point), when climate was generally cool and ice sheets were more or less at their largest, based on ice core and sea level records (Mix et al., 2001). It represents the time of maximum terrestrial

1 ice volume. More recently, the last sea level lowstand has been found to have occurred either
2 around 26 ka (Peltier and Fairbanks, 2006) or 21 ka (Lambeck et al., 2014) with relatively stable
3 (low) sea level between those dates. Nearly all ice sheets were at or close to their maximum
4 extent between 26 ka and 19 ka (Clark et al., 2009).

5 During the LGM, global annual mean surface temperatures are estimated to have been around
6 4.0 ± 0.8 °C colder than today (Annan and Hargreaves, 2013). The Earth began warming
7 towards its present state from around 19 ka (Fig. 1h; Buizert et al., 2014; Jouzel et al., 2007),
8 as summer insolation at northern high latitudes and global atmospheric greenhouse gas
9 concentrations gradually increased (Fig. 1c-f; Bereiter et al., 2015; Berger, 1978; Loulergue et
10 al., 2008; Marcott et al., 2014). By 9 ka, although the northern ice sheets had not quite retreated
11 (or disappeared) to their present day configuration, most of the Northern Hemisphere
12 deglaciation had taken place (Clark et al., 2012; Lambeck et al., 2014; Peltier et al., 2015;
13 Tarasov et al., 2012; Figures 1g and 2), with both surface air temperatures (Fig. 1h-i) and
14 atmospheric greenhouse gases (Fig. 1d-f) approaching present day values. However, much of
15 Antarctica remained heavily glaciated well into the Holocene, with the majority of its [deglacial](#)
16 [ice melting loss taking place](#) between 12 and 6 ka (Argus et al., 2014; Briggs et al., 2014;
17 Mackintosh et al., 2014). Antarctica's total contribution to post-glacial eustatic sea level is
18 poorly constrained, but recent studies have not supported LGM contributions greater than about
19 15 m eustatic sea level equivalent (Bentley et al., 2014; Briggs et al., 2014; Golledge et al.,
20 2013; Mackintosh et al., 2011; Philippon et al., 2006; Whitehouse et al., 2012), emphasising
21 the dominance of North American and Eurasian Ice Sheet dynamics in the global sea level
22 record during the last deglaciation (Argus et al., 2014; Lambeck et al., 2014; Peltier et al., 2015).
23 It should be noted that there is some controversy over whether deglacial ice sheet
24 reconstructions close the global sea level budget (Clark and Tarasov, 2014), with a potential
25 LGM shortfall of 'missing ice'.

26 The last deglaciation is not only an interesting case study for understanding multi-millennial
27 scale processes of deglaciation, but also provides the opportunity to study shorter and more
28 dramatic climate changes. Superimposed over the gradual warming trend (EPICA Community
29 Members, 2004; Jouzel et al., 2007; Petit et al., 1999; Stenni et al., 2011) are several abrupt
30 climate transitions lasting from a few years to a few centuries (examples of which are given

below) and it remains a challenge to reconstruct or understand the chain of events surrounding these instances of rapid cooling and warming.

Heinrich Event 1 (approx. 16.8 ka; Hemming, 2004) occurred during the relatively cool Northern Hemisphere Heinrich Stadial 1 (~18-14.7 ka). It was characterised by the release of a vast number of icebergs from the North American and Eurasian ice sheets into the open North Atlantic, where they melted. The existence of these iceberg ‘armadas’ is evidenced by a high proportion of ice rafted debris in North Atlantic sediments between 40° N and 55° N, predominantly of Laurentide (Hudson Strait) provenance (Hemming, 2004 and references therein). There are several competing theories for the cause of Heinrich Event 1. There is a substantial body of evidence ~~to suggest~~ suggesting that it occurred during or was precursory to a period of Atlantic Meridional Overturning Circulation (AMOC) slow down (e.g. Hall et al., 2006; Hemming, 2004; McManus et al., 2004) and weak North Atlantic Deep Water (NADW) formation (e.g. Keigwin and Boyle, 2008; Roberts et al., 2010) under a relatively cold, Northern Hemisphere surface climate (Shakun et al., 2012). Even though the interpretation of a cause and effect link between Heinrich Event 1 and the diminished strength of the AMOC remains rather compelling (e.g. Kageyama et al., 2013), it is increasingly being suggested that the melting icebergs might not have caused the recorded AMOC slow down, but may have provided a positive feedback to amplify or prolong AMOC weakening and widespread North Atlantic cooling (e.g. Álvarez-Solas et al., 2011; Barker et al., 2015), ~~whilst also causing mid-latitude Atlantic sea surface warming through northward expansion of the subtropical gyre.~~

During the subsequent 14.2-14.7 ka interval, Northern Hemisphere temperatures are seen to have risen by as much as 14.4 ± 1.9 °C in just a few decades (Buizert et al., 2014; Goujon et al., 2003; Kindler et al., 2014; Lea et al., 2003; Severinghaus and Brook, 1999), with a dramatic shift in some components of Greenland climate taking place in as little as one to three years (Steffensen et al., 2008). This abrupt event is termed the *Bølling Warming* or *Bølling Transition* (Severinghaus and Brook, 1999). At roughly the same time (~14.6 ka), there was a rapid jump in global sea level of 12-22 metres in around 350 years or less, known as *Meltwater Pulse 1a* (MWP1a; Deschamps et al., 2012). It is not known exactly which ice mass(es) contributed this 40 mm yr⁻¹ (or greater) flux of water to the oceans (e.g. Lambeck et al., 2014; Peltier, 2005). ~~Some older~~ studies have mainly attributed it to a southern source (Bassett et al., 2005, 2007; Carlson, 2009; Clark et al., 1996, 2002; Weaver et al., 2003), whereas more recent work has

suggested that at most, less than 4.3 metres eustatic sea level equivalent of meltwater could have come from Antarctica (Argus et al., 2014; Bentley et al., 2010, 2014; Briggs et al., 2014; Golledge et al., 2012, 2013, 2014; Licht, 2004; Mackintosh et al., 2011, 2014; Whitehouse et al., 2012) and that Northern Hemisphere ice was the primary contributor (Aharon, 2006; Gregoire et al., 2012; Keigwin et al., 1991; Marshall and Clarke, 1999; Peltier, 2005; Tarasov et al., 2012; Tarasov and Peltier, 2005). Exactly how the Bølling Warming and MWP1a are linked, or what triggered either, remains uncertain.

Ice core records of δD indicate that from around 14.5 ka to 12.8 ka, the general trend of increasing Southern Hemisphere warming, temporarily stalled (Jouzel et al., 2007; ice core chronology from Veres et al., 2013) for a period known as the *Antarctic Cold Reversal* (Jouzel et al., 1995). Southern Hemisphere cooling is thought to have been relatively widespread, extending from the South Pole to the southern mid-latitudes, with glacial readvance (or stall in glacial retreat) recorded to have peaked 13.0-14.2 ka in Patagonia (García et al., 2012; Kaplan et al., 2011; Strelin et al., 2011) and ~13.0 ka in New Zealand (Putnam et al., 2010; Rother et al., 2014). There are several hypotheses for the cause of the Antarctic Cold Reversal. For example, some have linked it to a change in ocean circulation induced by the delivery of Antarctic ice melt to the Southern Ocean (Menviel et al., 2010, 2011), or possibly as a bipolar response to AMOC recovery and Northern Hemisphere warming during the Bølling Warming (Menviel et al., 2011; Stocker, 1998). Using a CMIP5 level coupled atmosphere-ocean model, Peltier and Vettoretti (2014) and Vettoretti and Peltier (2015) have recently shown that ice core inferred southern hemisphere cooling and northern hemisphere warming could have been caused by a nonlinear salt oscillator mechanism. Others have argued that a change in Southern Hemisphere winds and ocean circulation is the explanation; for example, a simultaneous northward migration of the southern Subtropical Front and northward expansion of cold water originating in the Southern Ocean (Putnam et al., 2010). The ongoing disagreement over the timing, duration and extent of the Antarctic Cold Reversal means that its cause is difficult to pin down.

The next event of particular interest is the *Younger Dryas cooling*, when Northern Hemisphere temperatures are thought to have dropped by several degrees at 12.8-11.7 ka and most prominently in high latitudes (Buizert et al., 2014; Heiri et al., 2007; Lea et al., 2003; Liu et al., 2012; Simonsen et al., 2011; Steffensen et al., 2008). The event presents a conceptual paradox;

1 the magnitude of the cooling is difficult to reconcile with rising atmospheric CO₂
2 (approximately +10 ppm compared to the earlier Bølling period ~ 14.5 ka; Bereiter et al., 2015)
3 and increasing boreal summer insolation (Berger and Loutre, 1991). It is possible that changes
4 in the atmospheric hydrological cycle, such as a shift in source moisture region, could be partly
5 responsible for the $\delta^{18}\text{O}$ signal, requiring a smaller temperature anomaly to match the records
6 (Liu et al., 2012). For the climate cooling itself, a rerouting of North American freshwater
7 discharge to the Arctic and/or Atlantic Oceans might have caused a reduction in NADW
8 formation (Broecker et al., 1989; Condrón and Winsor, 2012; Tarasov and Peltier, 2005).
9 Simulating this period within the context of the preceding climate evolution, could be key to
10 understanding exactly what the surface climate and deep ocean changes were during the
11 Younger Dryas, and how these relate to contemporaneous proxy records (e.g. Buizert et al.,
12 2014).

13 In this description, we have sought to capture some of the last deglaciation's main climatic
14 events, but there are others that could shape the focus of further study in the working group.
15 For example, early on in the period there is evidence of around 10 m sea level rise taking place
16 in 500-800 years around 20-19 ka (Clark et al., 2004; Clark and Mix, 2002; De Deckker and
17 Yokoyama, 2009; Yokoyama et al., 2001a, 2001b). Whilst the event itself remains somewhat
18 controversial (Cabioch et al., 2003; Hanebuth et al., 2000, 2009; Peltier and Fairbanks, 2006;
19 Shennan and Milne, 2003), it could be the expression of accelerating deglacial ice melt
20 following the Last Glacial Maximum. More recently, the Barbados record of relative sea level
21 history indicates that following the Younger Dryas cooling episode, there may have been
22 another meltwater pulse (Fairbanks, 1989; Peltier and Fairbanks, 2006), referred to as
23 Meltwater Pulse 1b. Significant debate surrounds the magnitude and timing of Meltwater Pulse
24 1b (Bard et al., 1996; Cabioch et al., 2003; Cutler et al., 2003; Edwards et al., 1993; Shennan,
25 1999; Stanford et al., 2011) and even its existence, because similar to the 19 ka event, it is not
26 seen in all sea level records spanning the interval (e.g. Bard et al., 1996, 2010; Hanebuth et al.,
27 2000). However, evidence of rapid Antarctic retreat around the time of the event could provide
28 a possible cause for this late deglacial rapid sea level rise (Argus et al., 2014).

1.2 Transient modelling of the last deglaciation

Transient modelling of the last deglaciation is valuable for examining dynamic and threshold behaviours (Braconnot et al., 2012) endemic to the Earth's non-stationary climate system, especially ice-ocean-atmosphere interactions. It is the best tool for reaching a comprehensive understanding of complex and interrelating climate processes with specific regard to chains of events.

Such simulations are useful for examining the effect of temporally varying climate forcings across the globe and in different environmental systems: what geographical patterns arise and how are they connected, how do these vary through time from seasonal to millennial time scales, and how long does it take before a change in forcing is manifested in a climate response? The spatial coherency of specific events can be investigated to identify processes for simultaneous change as well as lead/lag mechanisms. For example, Roche et al. (2011) investigated patterns of spatial variability in the deglaciation as caused by long-term changes in orbital parameters, atmospheric greenhouse gas concentrations, and ice sheet extent/topography. The results indicated a simultaneous onset of hemispheric warming in the North and South, showing that obliquity forcing was the main driver of the early deglacial warming. In the same investigation, it was found that sea-ice covered regions were the first parts of the world to exhibit significant rises in temperature, implying that a better knowledge of sea-ice evolution could be key to fully understanding the trigger for widespread deglaciation and warming feedbacks. A further example of the insights available into lead-lag relationships provided by long, transient climate simulations under glacial boundary conditions is provided by the previously referenced Dansgaard-Oeschger oscillation-related analyses of Peltier and Vettoretti (2014) and Vettoretti and Peltier (2015), which appear to mimic the Heinrich Stadial 1 to Bølling transition.

Through comparison to geological timeseries data, transient simulations enable the 'fingerprinting' of specific climate processes to find out what mechanisms [in the model] can cause recorded climate signals. Comparing complex, global-scale models to combined geological records can provide multiple 'fingerprints' in different variables from different archives and in different locations to help narrow down plausible scenarios. For example, Menviel et al. (2011) ran a suite of simulations, varying oceanic meltwater fluxes through the last deglaciation in order to identify which freshwater-forcing scenarios reproduce the Atlantic

1 Ocean circulation state implied by sedimentary records of AMOC strength/depth and
2 ventilation age (Gherardi et al., 2005; McManus et al., 2004 with ages shifted as per Alley,
3 2000; Thornalley et al., 2011) as well as the Northern Hemisphere surface climate (Alley, 2000;
4 Bard, 2002; Bard et al., 2000; Heiri et al., 2007; Lea et al., 2003; Martrat et al., 2004, 2007). It
5 was argued that such climate simulations could be used to improve constraints on the timing,
6 duration, magnitude, and location of meltwater inputs to the global ocean.

7 Liu et al. (e.g. 2009) used climate ‘fingerprinting’ to identify possible mechanisms for the
8 abrupt Bølling Warming Event, finding that in their model, a forced cessation of freshwater
9 inputs to the North Atlantic (representing ice sheet melt) superimposed on a steady increase in
10 atmospheric CO₂ caused an abrupt resumption in the strength of the AMOC (almost matching
11 a record produced by McManus et al., 2004). This in turn induced a rapid warming in Northern
12 Hemisphere surface climate (close to records from Bard et al., 2000; Cuffey and Clow, 1997;
13 and Waelbroeck et al., 1998) and an increase in tropical rainfall over the Cariaco Basin
14 (comparable to Lea et al., 2003), whilst Antarctic surface temperatures remained relatively
15 stable (similar to Jouzel et al., 2007). Using a suite of simulations from the same model, Otto-
16 Bliesner et al. (2014) went on to suggest that a combination of rapid strengthening of NADW
17 seen by Liu et al. (e.g. 2009) and rising greenhouse gas concentrations was responsible for
18 increased African humidity around 14.7 ka, matching the model output to a range of regional
19 climate proxies (including deMenocal et al., 2000; Tierney et al., 2008; Tjallingii et al., 2008;
20 Verschuren et al., 2009; Weijers et al., 2007).

21 Thus, climate proxy fingerprinting can be useful for understanding the spatial coherency of
22 climatic changes and their underlying mechanisms. However, correlation between model and
23 geological data does not guarantee that the correct processes have been simulated; there is
24 always the problem of *equifinality*, whereby the same end state can be reached by multiple
25 means. In a process sense, this may be particularly uncertain when a model does not reproduce
26 the full chain of events that led to a distinguishable climatic signal. For example, mechanisms
27 for many of the major changes in oceanic freshwater inputs proposed by Liu et al. (2009) and
28 Meniel et al. (2011) have not yet been directly simulated (e.g. by dynamic ice sheet models).
29 In both studies, they are imposed as model boundary conditions. Further simulations with
30 different forcing scenarios and from a range of models would help to address such uncertainties.

Transient simulations of the last deglaciation also provide necessary boundary conditions for modelling a variety of Earth System components that may not be interactively coupled to the climate model being used. For example, Gregoire et al. (2015) drove a dynamic ice sheet model with climate data produced by a similar set of simulations to Roche et al. (2011). Using a low resolution GCM, individual climate forcings – including orbit, greenhouse gases, and meltwater fluxes – were isolated so that their relative contribution to melting the modelled North American ice sheets could be examined. The work concluded that the last deglaciation was primarily driven by changes in Northern Hemisphere insolation, causing around 60% of the North American Ice Sheet melt, whilst increasing CO₂ levels were responsible for most of the remaining changes (Gregoire et al., 2015). The sufficiency of these two forcings for North American glaciation/deglaciation had previously also been identified with fully coupled glaciological and energy balance climate models (Tarasov and Peltier, 1997). Gregoire et al (2012) were also able to highlight a ~~possible~~ ‘saddle-collapse’ mechanism, whereby gradual warming trends could result in abrupt ice sheet melting events, ~~such as MWP1a and the 8.2 kyr Event, when a threshold in ice mass balance was crossed. The opening of the ice-free corridor between the Cordilleran and Laurentide ice sheets has long been built into the ICE-NG, Tarasov and Peltier and Tarasov et al. sequence of models as geological inferences indicate that it, which could have occurred around the same time as during MWP1a- and the 8.2 kyr event..~~

A further example is given by Liu et al. (2012), who carried out an asynchronous (or ‘offline’) coupling between simulated sea surface temperatures and an isotope-enabled atmospheric model to investigate the Younger Dryas cooling event (~12 ka). The results revised the presupposed Greenland temperatures at this time by 5 °C, demonstrating that changes in moisture source must be an important consideration for the robust interpretation of Greenland ice core $\delta^{18}\text{O}$ records and our understanding of high-latitude climate sensitivity. More recently, the same methodology was applied to understanding Chinese cave records of the East Asian Summer Monsoon 21-0 ka (Liu et al., 2014), not only to better interpret what the speleothem $\delta^{18}\text{O}$ tells us about regional hydroclimate variability, but also to understand the wider teleconnections controlling those patterns.

In addition, there are now transient simulations of the last deglaciation from climate models that have been interactively coupled with dynamic ice sheet models (Bonelli et al., 2009; Heinemann et al., 2014) and isotope systems (Caley et al., 2014). Furthermore, a fast Earth

System Model of Intermediate Complexity (EMIC) that includes an interactive ice sheet model has been used to look at Earth System dynamics (the role of orbital cycles, aeolian dust, subglacial regolith properties, the carbon cycle, and atmospheric trace gases) on much longer, glacial-interglacial timescales >120 ka and encompassing the last deglaciation (Bauer and Ganopolski, 2014; Brovkin et al., 2012; Ganopolski et al., 2010; Ganopolski and Calov, 2011). However, the older, uncoupled climate-ice sheet model approach discussed above remains useful because it enables a wider suite of models to be employed than would otherwise be feasible due to limited computational efficiency (e.g. of state-of-the-art, high resolution/complexity models) or software engineering capability. It may also allow for the same Earth System component model (e.g. of ice sheets or $\delta^{18}\text{O}$) to be driven by multiple climate models, in order to examine the range of responses and assess [climate] model performance.

With sufficient computational power to make long simulations of the last deglaciation a feasible undertaking, it is timely to coordinate new efforts to ensure that a framework exists to (i) utilise the cutting edge science in climate modelling and palaeoclimate reconstruction, and (ii) robustly intercompare simulations run with different models by different groups and palaeoclimatic data.

1.3 Establishing a new PMIP working group

For [more than](#) twenty years, the Paleoclimate Modeling Intercomparison Project (PMIP) has been internationally coordinating multi-model simulations with complex climate models in order to evaluate model performance and better understand [past] climate changes (Braconnot et al., 2007, 2012; PMIP website, 2007). Currently entering its fourth phase, PMIP is a growing organisation that continues to contribute towards other coordinated efforts to understand present day climate change; including the Coupled Model Intercomparison Project (Taylor et al., 2011a, CMIP; e.g. 2011b) and the Intergovernmental Panel on Climate Change's (IPCC) Assessment Reports (e.g. the Fifth Assessment Report; Flato et al., 2013; Masson-Delmotte et al., 2013). It encompasses a broad range of models, from very fast, lower resolution EMICS, through a range of coupled GCMs to the latest generation of higher resolution and complexity Earth System Models. Thus, the main challenges for the fourth Phase of PMIP include: designing experiments that are suitable for all of its participants; addressing sufficiently fundamental questions to be of interest to the EMIC community; defining adequately focused scope for the feasible

participation of the latest generation of ESMs; and prescribing flexible model setups that can be implemented in this range of models, whilst maintaining the ability to robustly compare results. [In addition, a continuing challenge for PMIP is to assemble suitable palaeoclimatic datasets for comparison to model results.](#)

One of the most recent working groups to be established in PMIP is the Last Deglaciation Working Group. With the aim of coordinating transient simulations of the last deglaciation, the challenge of including the full range of PMIP models is at the forefront of our experiment design. The experiment will be partitioned into three phases (Fig. 1b and Sect. 4), which will form milestones for managing its long duration (12 thousand years) as well as for scheduling any shorter, alternative simulations to the Core.

The aim of this paper is to outline the model setup for the transient Core [simulation of experiment for](#) the last deglaciation, specifically for the sub-period of 21-9 ka. Prescribed boundary conditions include orbital parameters, atmospheric trace gases and ice sheets. In association with the ice sheet reconstructions, we also provide bathymetric, orographic and land-sea mask evolution, [as well as make recommendations for freshwater forcing \(or global ocean salinity changes\) through the period.](#)

1.4 Approach

One of the roles of PMIP has been to systematically study the ability of climate models to retrodict different past climates for which there are ‘observational’ data from geological archives (e.g. Braconnot et al., 2000, 2007, 2012; Haywood et al., 2010; Joussaume et al., 1999; Kageyama et al., 2006; Kohfeld and Harrison, 2000; Masson-Delmotte et al., 2006; Otto-Bliesner et al., 2009; Weber et al., 2007). In this vein, many palaeoclimate model intercomparison projects have been designed to facilitate the robust comparison of results from the same ‘experiment’ (i.e. simulation set) across a range of different models, usually taking a prescriptive approach to model setup to ensure that any differences observed in the results are attributable to differences in model structure and not to differences in chosen ‘boundary conditions’ and climate forcings. However, as Schmidt et al. (2011) point out, the choice of one particular configuration from a range of plausible boundary conditions and forcings is often arbitrary and does not account for uncertainties in the data used for developing the forcings/boundary conditions. Moreover, in designing the PMIP last deglaciation experiment,

1 we have attempted to strike a balance between establishing a framework within which to assess
2 model differences and performance, and taking the opportunity to utilise the full range of PMIP
3 climate models (Earth System, General Circulation and Intermediate Complexity) to examine
4 uncertainties in deglacial forcings, trigger-mechanisms and dynamic feedbacks. In short, when
5 we do not precisely know the climate forcing for an event, or the temporal evolution of model
6 boundary conditions, it is more efficient to compare the results from models that use different
7 forcings with geological and palaeoclimatic data than to run one scenario with all models and
8 all scenarios with all models. The aim is to use the results of the comparison to narrow down
9 the range of uncertainty in the forcings/boundary conditions and reach a better understanding
10 of underlying climate mechanisms.

11 Consequently, forcings/boundary conditions that are relatively well established (atmospheric
12 trace gases and orbital parameters) are tightly constrained in the Core experiment design. Others
13 are given with multiple precisely described possibilities to choose from (ice sheet
14 reconstructions) and the remainder (e.g. freshwater/salinity, aerosols and vegetation) are left to
15 the discretion of individual participants, ~~although we recommend~~. Recommendations will be
16 made for the latter grouping of forcings/boundary conditions; for example, freshwater/global
17 salinity fluxes that are consistent with the provided ice sheet evolutions, and the use of
18 preindustrial aerosol and/or vegetation values when they are not model prognostics; but a
19 flexible approach is advantageous not only scientifically (i.e. for examining the climatic
20 response to uncertain forcings, see above), but also practically (for accommodating the wide
21 range of participating models). Further to this, it will be left to the expert user to decide how
22 often to make manual updates to those boundary conditions that cannot evolve automatically in
23 the model, such as bathymetry, orography and land sea mask. This is also necessary because of
24 the specific technical and resource requirements associated with setting up and running each
25 participant model.

26 In addition to the Core, we will ~~also~~ coordinate ~~additional~~ a series of experiments that are
27 designed to:

- 28 (i) explore uncertainties in the boundary conditions and climate forcings,
- 29 (ii) test specific hypotheses for mechanisms of climate change and to explain individual
30 events.

(iii) focus on shorter time periods (for example, abrupt events) and thus include computationally expensive models for which a twelve thousand year simulation is unfeasible.

These optional simulations will be referred to as *focussed* experiments, and participants are encouraged to contribute towards the design and coordination of these simulations within the working group ([dedicated Wiki page to coordinate these here: https://wiki.lsce.ipsl.fr/pmip3/doku.php/pmip3:wg:degla:index](https://wiki.lsce.ipsl.fr/pmip3/doku.php/pmip3:wg:degla:index)).

The start date for the experiment has been chosen to be in line with PMIP's historical definition of the LGM; 21 ka (Abe-Ouchi et al., 2015; e.g. Braconnot et al., 2000; Kohfeld and Harrison, 2000). However, we are aware that some groups may prefer to begin their simulations from the earlier date of 26 ka (around the last sea level lowstand; Clark et al., 2009; Lambeck et al., 2014; Peltier and Fairbanks, 2006) and both orbital and atmospheric trace gas parameters will be provided from this earlier date. Although the working group's focus will at least initially be 21-9 ka, boundary conditions for the Core ~~simulation~~[simulations](#) will be provided from 21 ka to the preindustrial (26 ka to the preindustrial for orbital insolation and trace gases).

The following is not meant to be an exhaustive review of climate forcing reconstructions through the last deglaciation. Instead, our intention is to consolidate the current knowledge in a practical experiment design for a range of climate models. Within this coordinated context, the aim is to explore the forcings and underlying feedback mechanisms for the rapid climate events that punctuated the gradual warming and deglaciation of the Earth.

The paper is structured so that Sect. 2 outlines the model boundary conditions and climate forcings for the Core ~~simulation~~[experiment](#). Section 3 presents how we will ensure the feasible participation of a range of climate models with different complexity and computational efficiency, as well as the plan to run additional, targeted, hypothesis- and sensitivity-led simulations. Section 4 discusses the three phases of the long Core experiment.

2 Core ~~simulation~~[experiment](#) (21 ka to 9 ka)

The Core ~~simulation~~[simulations](#) of the last deglaciation will focus on the period from 21 ka to 9 ka, although there will also be the option to spin up the simulation with time-evolving orbital and trace gas parameters from 26 ka and all boundary conditions will be available from 21 ka to the preindustrial. Recommendations for the initialisation state at 21 ka are summarised

in Table 1 and described below (Sect. 2.1). Prescribed boundary conditions include insolation via the Earth's astronomical parameters (Sect. 2.2), atmospheric trace gases (Sect. 2.3), ice sheets (Sect. 2.4), meltwater fluxes (Sect. 2.5), and orography/bathymetry (Sect. 2.6), as summarised in Table 2. Boundary condition data for the Core [simulationexperiment](#) are provided on the PMIP wiki; <https://wiki.lsce.ipsl.fr/pmip3/doku.php/pmip3:wg:degla:bc:core> (PMIP Last Deglaciation Working Group, 2015).

2.1 Last Glacial Maximum spinup

There is a choice of two possibilities for starting the last deglaciation Core [simulations](#). Either the simulation should be initialised from the end of a spun-up, PMIP-compliant LGM (21 ka) simulation, or a simulation with transient orbital and trace gas forcing should be run from an earlier time period (orbital and trace gas parameters will be provided from 26 ka onwards). Whichever method is applied, we require that it is comprehensively documented along with information on the model's state of spinup at 21 ka (e.g. timeseries of surface climates, maximum strength of the North Atlantic Meridional Overturning Circulation stream function, net radiation at the top of the atmosphere [etc.](#)).

2.1.1 Equilibrium-type spinup (21 ka)

For setting up an equilibrium-type spinup, please make sure to use the following constraints, which may differ from other PMIP 21 ka simulation protocols:

- Insolation should be set so that eccentricity is 0.018994, obliquity is 22.949°, perihelion-180° is 114.42°, the date of the vernal equinox is 21st March at noon, and the solar constant is the same as for the preindustrial (e.g. 1365 W m⁻², as in the PMIP3-CMIP5 preindustrial experiment). These are consistent with previous PMIP LGM boundary conditions (PMIP LGM Working Group, 2010).
- Prescribed atmospheric trace gases should be as follows: CO₂ at [188190](#) ppm, CH₄ at 375 ppb, N₂O at 200 ppb (Fig. 3), with CFCs at 0 and O₃ at the PMIP3-CMIP5 preindustrial value (e.g. 10 DU). This is to be compatible with the time-evolving boundary conditions for the Core [simulations](#) (Sect. 2.3). Note that the LGM atmospheric CO₂ and CH₄ concentrations have changed slightly from earlier LGM experiments (e.g. PMIP3, which used 185 ppm and 350 ppb, respectively;

PMIP LGM Working Group, 2010). However, N₂O remains at 200ppb, which is more representative of the longer glacial period than the 187 ppb concentration recorded at 21 ka (Fig. 3c). These updates are in line with the latest ice core age model, (AICC2012; Veres et al., 2013) and records (Bereiter et al., 2015; Schilt et al., 2010), ~~which is,~~ which are also used for the transient forcings described below (Sect. 2.3).

- Prescribed ice sheets should use either the GLAC-1D or ICE-6G_C reconstruction at 21 ka (see Sect. 2.4). The associated topography and coastlines should be used as per the chosen ice sheet reconstruction. Beyond maintaining consistency with the coastlines, it is optional whether or not to implement the associated bathymetry. and participants should adapt the bathymetry according to their model's capabilities (for example, depending on whether the spatial resolution allows for it or makes this a useful adaptation). These data will be provided with the ice sheet reconstructions. Whichever ice sheet reconstruction is chosen for the LGM spinup should be carried through to the Core transient simulation.
- Global ocean salinity should be +1 psu, compared to preindustrial, to account for the increased terrestrial ice mass at the LGM (PMIP LGM Working Group, 2015).
- Any other boundary conditions should be set to be consistent with the Core transient simulation to follow (Sect. 2.2-2.7).

On the freshwater budget, PMIP advises groups to ‘carefully check the fresh water budget in their LGM experiments in order to avoid unnecessary drifts of the ocean salinity. It can be necessary to route the snow which has fallen in excess on the ice sheets to the ocean. Given the change in coastlines, it is also sometimes necessary to relocate the large river estuaries on the coast’ (PMIP LGM Working Group, 2015). Tarasov and Peltier (2006) ~~provides~~provide a glaciological example of the possible re-routings for North America. As they become available, routing maps for the Last Glacial Maximum continents will be provided on the last deglaciation PMIP Wiki (address above).

The integration time required for spinning up the LGM climate state should be decided on a case-by-case basis by the user. Groups may choose to initialise their equilibrium-type simulation from other PMIP LGM runs. However, ~~please be careful-~~caution is advised. Some

of the boundary conditions for ~~the PMIP4-CMIP6 (not finalised at the time of writing)~~ and previous PMIP LGM simulations are different to the setup outlined here, specifically in terms of ice sheets and trace gases concentrations, and therefore need to be adapted to match these requirements. ~~Please also~~ The protocol for the PMIP4-CMIP6 (being finalised at the time of writing) is currently compatible with the LGM spin-up described here. Therefore, provided that either the ICE-6G C or GLAC-1D ice sheet reconstruction is used for both the LGM spin-up and transient run, the PMIP4-CMIP6 LGM simulation can be used to initialise transient simulations of the last deglaciation without alteration. Please provide timeseries data for the diagnosis of model [dis]equilibrium at 21 ka (introduction to Sect. 2.1).

2.1.2 Transient orbital and trace gas parameters (26-21 ka)

If this is the preferred option to initialise the Core, it is recommended that the simulation is setup as per Sect. 2.1.1, but with time-evolving orbital and trace gas parameters instead of fixed ones. Specifically for orbit, the eccentricity, obliquity, perihelion–180° and date of the vernal equinox values listed above should be replaced with their transient equivalents, as per Berger (1978). For the atmospheric trace gases, carbon dioxide, methane and nitrous oxide values should be replaced with the transient equivalents provided on the PMIP Wiki (<https://wiki.lsce.ipsl.fr/pmip3/doku.php/pmip3:wg:degla:bc:core>) and according to LüthiBereiter et al. (2015), Loulergue et al. (2008) and Schilt et al. (2010), respectively, on the AICC2012 chronology (Veres et al., 2013); Fig. 3.

In this case, all other boundary conditions should remain fixed in line with the LGM equilibrium-type experiment design until 21 ka, when the fully transient Core ~~simulation begins~~ simulations begin. This transient spin-up can be initialised from a spun-up previous LGM, cold ocean, preindustrial, or observed present day ocean simulation.

2.2 Insolation (21-9 ka)

As per Sect. 2.1, the solar constant should be fixed to the established preindustrial conditions (e.g. 1365 W m⁻²) throughout the run, which is the PMIP preindustrial experiment setup (PMIP LGM Working Group, 2015). However, the orbital parameters should be time-evolving through the deglaciation to follow Berger (1978); e.g. Fig. 1c.

2.3 Atmospheric trace gases (21-9 ka)

For the deglaciation, CFCs should be fixed at 0, and O₃ should be set to PMIP3-CMIP5 preindustrial values (e.g. 10 DU), as used for the LGM. When a model is not running with dynamic atmospheric chemistry, the remaining trace gases should be time-evolving, with CO₂ following LüthiBereiter et al. (2015), CH₄ following Loulergue et al. (2008) and N₂O following Schilt et al. (2010), all adjusted to the AICC2012 chronology (Veres et al., 2013); Fig. 1d-f.

The atmospheric CO₂ concentrations provided by Bereiter et al. (2015) is a composite dataset, combining previous Antarctic ice core records and composites (for the period 26-0 ka: Ahn and Brook, 2014; Lüthi et al., 2008; MacFarling Meure et al., 2006; Marcott et al., 2014; Rubino et al., 2013; Siegenthaler et al., 2005) on the AICC2012 timescale of Veres et al. (2013)~~Temporally higher resolution CO₂ data from the West Antarctic Ice Sheet Divide has been provided by Marcott et al., spanning 23-9 ka ('WDC' on Fig. 3a). However, the newer data are consistently offset from other Antarctic ice core data by ~4 ppm and the cause for this remains unresolved. Furthermore, although the data encompasses the last deglaciation (and the period we are focussing on; 21-9 ka), it would not be easily spliced into a longer record (e.g. for groups wishing to run their simulations through to the present day). This is why the higher resolution data will not be used for the Core, reverting to the older record from Lüthi et al.. However, it to produce a high resolution record that is consistent with the other, lower resolution trace gas records used in this experiment (CH₄ and N₂O as discussed above). Groups are free to decide on the temporal resolution of trace gas model inputs based on these records and if lower resolution is employed, the method used to smooth or create a spline through the data should be fully documented. Exploring the influence of CO₂ resolution on the climate system~~ may form the basis of a coordinated additional simulation, which will be optional for participant groups. ~~Other sensitivity type simulations could also be coordinated to assess the influence of timing in the CO₂ records on climate and ice sheet evolution, addressing age model uncertainty.~~ The details of the setup for such *focussed* simulations (also discussed in Sect. 3) will be discussed and determined at a later date.

It is noted that the N₂O value from Schilt et al. (2010) and Veres et al. (2013) does not match the previously defined LGM N₂O concentration (Sect. 2.1.1); 187 ppb compared to 200 ppb (Fig. 3c). This is because the N₂O record is highly variable during the last glacial lowstand (26-21 ka), with a range of ~33 ppb (183-216 ppb) and a mean of 201 ppb. Thus 200 ppb seems a

reasonably representative N₂O concentration for the spinup phase of the simulation, although the Core ~~simulation~~[simulations](#) will start with the more chronologically accurate value of 187 ppb.

2.4 Ice sheet reconstructions (21-9 ka)

For the Core experiment, ice sheet extent and topography should be prescribed from one of two possible reconstructions: ICE-6G_C (Fig. 2a and Fig. 4a) and GLAC-1D (Fig. 2b and Fig. 4b).

The ICE-6G_C reconstruction is fully published (Argus et al., 2014; Peltier et al., 2015), and the reader is directed to this literature for further information. The GLAC-1D reconstruction is combined from different sources (Briggs et al., 2014; Tarasov et al., 2012; Tarasov and Peltier, 2002) and whilst it is mostly published, there are some new components; therefore, a short description follows. The Eurasian and North American components are from Bayesian calibrations of a glaciological model (Tarasov et al., 2012; this study), the Antarctic component is from a scored ensemble of 3344 glaciological model runs (Briggs et al., 2014) and the Greenland component is the hand-tuned glaciological model of Tarasov and Peltier (2002) [updated to the GICC05 age chronology](#) (Rasmussen et al., 2006). All four of the GLAC-1D ice sheet components employ dynamical ice sheet models that have been constrained with relative sea level data. Where available, they have also been constrained by geologically-inferred deglacial ice margin chronologies, pro-glacial lake levels, ice core temperature profiles, present-day vertical velocities, past ice thickness, and present day ice configuration. Details of exactly how these constraints were derived and applied are given in the relevant references above. The four components (North American, Eurasia, Antarctica and Greenland) were combined under Glacial Isostatic Adjustment (GIA) post-processing for a near-gravitationally self-consistent solution (Tarasov and Peltier, 2004), which was tested against complete Glacial Isostatic Adjustment solutions ~~(Tarasov, pers. comm. 2014)~~. The topography in the global combined solution was adjusted in Patagonia and Iceland following ICE-5G (Peltier, 2004), but the changes in these ice caps are not reflected in the ice mask.

Both datasets include ice extent and topography at intervals of 1,000 years or less through the deglaciation. ~~I.e.~~ [Specifically, the ICE-6G_C reconstruction is provided at 1,000-year intervals for the period spanning 26-21 ka and 500-year intervals for 21-0 ka. For GLAC-1D, the data](#)

are at 100-year intervals 21-0 ka. In both reconstructions, ice extent is provided as a fractional ice mask for ICE-6G_C and a binary ice mask in GLAC-1D.

The two reconstructions incorporate similar constraints for North American ice sheet extent (i.e. Dyke, 2004). For Eurasia, ICE-6G_C follows the ice extent provided by Gyllencreutz et al. (2007), whereas GLAC-1D uses data from Hughes et al. (2015). The reconstructions only differ slightly in their ice extent evolution (Figures 2 and 4), for example the Barents Sea deglaciates earlier in GLAC-1D than in ICE-6G_C (Fig. 2). The main differences between the reconstructions are in the shape and volume of individual ice sheets. In particular, the North American Ice Sheet reaches an elevation of 4000 m in ICE-6G_C, but is only 3500 m high in GLAC-1D. Similarly, the shape and thickness of the Barents Sea Ice Sheet are not the same in the two reconstructions. The ICE-6G_C dataset is ~~been~~ provided at both 1 degree horizontal resolution and 10 minute horizontal resolution, GLAC-1D is provided at 1 degree (longitude) \times 0.5 degree (latitude) horizontal resolution.

Ice surface elevation (topography) should be implemented as an anomaly from present day topography and added to the model's present day topography after regridding onto the model resolution, following the previous LGM experimental protocol (PMIP LGM Working Group, 2010, 2015). Land surface properties will need to be adjusted for changes in ice extent. Where ice retreats, land surface should be initialised as bare soil if a dynamic vegetation model is used, otherwise use prescribed vegetation (see Sect. 2.7) with appropriate consideration of soil characteristics. Where ice is replaced by ocean, it is advised to follow the procedure for changing coastlines described in Sect. 2.7. Inland lakes can be prescribed based on the ice sheet and topography reconstructions, but this is not compulsory. It is also optional whether to include changes in river routing basins (i.e. catchments) and outlets, which can either be calculated from the provided topography and land-sea mask data (see Sect. ~~2.6~~2.6), or can be manually set to follow routing maps, which will be provided on the last deglaciation PMIP Wiki.

Groups are free to choose how often to update ice extent and elevation. This could be done at regular intervals (e.g. the sub-1000 year time slices provided) or at specific times during the deglaciation, as was done in the TraCE-21 ka experiment (Liu et al., 2009). Changes in ice extent can have a large impact on climate through ice albedo changes and feedbacks. We thus recommend that when possible, ice sheets are not updated at times of abrupt regional or global climate change, particularly the events that the working group will focus on, as this could

1 artificially introduce stepped shifts in climate. Groups are also advised to consider that ice sheet
2 associated boundary conditions (ice extent and elevation, land-sea mask, bathymetry) may need
3 to be updated more often at times of rapid ice retreat. The timing and way in which land ice
4 changes are implemented must be documented.

5 Alternative ice sheet reconstructions or simulations can be used to test the sensitivity of climate
6 to this boundary condition. Simulations with coupled ice sheet-climate models are also
7 welcomed. Although these will not form part of the Core, for which ICE-6G_C or GLAC-1D
8 should be used, they will be coordinated as important supplementary *focussed* simulations.

9 **2.5 Ice meltwater**

10 The Core ~~simulation will~~experiment protocol is flexible on whether or not to include ~~any~~
11 prescribed ice melt (i.e. freshwater fluxes) delivered from the ice sheets to the ocean. ~~This may~~
12 ~~seem controversial given the levels and how to do it. It is recommended to run at least one~~
13 version of ~~terrestrial~~the Core experiment with ice-sheet melt ~~and included, since around 110 m~~
14 of ice-volume equivalent sea-level ~~rise known is thought~~ to have ~~taken place during this~~
15 period~~melted 26-9 ka~~ (e.g. Lambeck et al., 2014) and considering the historical importance
16 attached to the influence of [de]glacial freshwater fluxes on climate (e.g. Broecker et al., 1989;
17 Condon and Winsor, 2012; Ganopolski and Rahmstorf, 2001; Liu et al., 2009; Rahmstorf,
18 1995, 1996; Teller et al., 2002; Thornalley et al., 2010; Weaver et al., 2003). However,
19 ~~considering the current uncertainty on exactly when and where ice melt entered the ocean~~
20 ~~during the last deglaciation (e.g. discussion of MWP1a in Sect. 1.1), this~~ it is ~~the best way~~also
21 important to ~~ensure that the Core experiment is based on robust geological data. Furthermore,~~
22 ~~there is an~~ note the ongoing debate over the ~~role of~~extent to which catastrophic freshwater
23 fluxes ~~in bringing~~brought about abrupt deglacial climate change ~~and~~; several alternative or
24 complementary mechanisms have been proposed (e.g. Adkins et al., 2005; Álvarez-Solas et al.,
25 2011; Barker et al., 2010, 2015; Broecker, 2003; Hall et al., 2006; Knorr and Lohmann, 2003,
26 2007; Roche et al., 2007; Rogerson et al., 2010; Thiagarajan et al., 2014). ~~In light of this, and~~
27 ~~because we are keen to see what the climate response to non-freshwater forced scenarios will~~
28 ~~be in the PMIP models, the decision has been made to have no prescribed freshwater fluxes in~~
29 ~~the Core simulation. This experiment is thus designed to constitute a reference for experiments~~
30 ~~in which fresh water fluxes will be introduced.~~

Moreover, a thorough investigation of the extent to which non-freshwater-forced climate evolution matches the geological records has merit in its own right; can abrupt deglacial changes be simulated without ice-meltwater, ~~as has been proposed (e.g. discussion above)?~~ To what extent can ‘observed’ patterns be attributed to better constrained forcings, such as atmospheric CO₂ and Earth’s orbit? ~~To complete the investigation, freshwater flux scenarios will be targeted by opt-in focussed simulations that test specific ice-melt hypotheses as well as instances where/when the Core falls short of the ‘observed’ patterns. For example, routing of ice-melt computed from GLAC-1D (Sect. 2.4) will be provided as a possible transient boundary condition.~~ It is for all of these reasons that a flexible protocol is required.

Freshwater forcing scenarios consistent with the ice sheet reconstructions and which hence conserve salinity throughout the deglacial experiment are provided in two formats (the ‘melt-’ scenarios described below). In addition, there is the option to run without any ice meltwater (‘no-melt’) to provide a robust reference for simulations that include uncertain meltwater fluxes. Thus, at least one Core simulation should be run using one of the following ice sheet meltwater scenarios:

melt-uniform: a globally uniform freshwater flux (or salinity target) through time, designed to conserve ocean salinity based on changing terrestrial ice mass. Fluxes consistent with the ice sheet reconstructions are provided.

melt-routed: a distributed routing that is consistent with the geographic evolution of the ice sheet reconstructions (GLAC-1D and ICE-6G C; Sect. 2.4) and gives the flux through time at individual meltwater river outlets along the coast. Again, versions of this scenario are provided.

no-melt: no ice meltwater is included in the core; neither a globally integrated ocean salinity target (*melt-uniform*) nor a distributed routing at the coastlines (*melt-routed*) is implemented. This is best implemented as a sensitivity-type experiment to account for model-specificness and meltwater flux uncertainty when also implementing *melt-* scenarios in accompanying versions of the Core simulation.

Multiple Core simulations exploring more than one of these scenarios are welcomed.

1 Data for the *melt-* scenarios will be available from the PMIP last deglaciation Wiki. The data
2 for *melt-uniform* are available at the time of writing (following the respective ice volume
3 changes from ICE-6G_C and GLAC1-D; Fig. 1g), data for *melt-routed* will be made available
4 as they are produced (anticipated by April/May 2016). These *melt-* scenarios represent a ‘best-
5 estimate’ approach to resolving the yet unknown geographically- and temporally-precise
6 freshwater fluxes of the last deglaciation, and they are also consistent with the ice sheet
7 reconstructions employed in the core. As such, they provide robust and justifiable boundary
8 conditions for simulations that will be assessed against palaeoclimate reconstructions.

9 However, participants do not have to use the [recommended] versions of *melt-uniform* or *melt-*
10 *routed* that are consistent with ICE-6G_C and GLAC-1D, and can instead use their own
11 scenarios to explore uncertainty in the ice sheet meltwater flux forcing. This is because the
12 working group aims to use the full suite of PMIP climate models to examine forcing/boundary
13 condition uncertainty (see discussion of model intercomparison project approaches in Sect.
14 1.4). Please note that in some ice melt (including *no-melt*) scenarios, global water budget may
15 not be balanced through time (as is also true for *no-melt*). Therefore, it is advised to also use at
16 least one scenario that falls within geological constraints (such as the ICE-6G_C or GLAC-1D
17 consistent scenarios for *melt-uniform* and *melt-routed*).

18 Regardless of which scenario is employed, it is important that meltwater fluxes are prescribed
19 as time-evolving model boundary conditions; rather than as step-wise adjustments at the same
20 time as the ice sheets are updated, for example. Unless they are intentional conditions of the
21 scenario, there should be no sudden jumps in the freshwater being applied. Furthermore, we
22 invite participants to upload the boundary condition data for other freshwater flux scenarios
23 along with appropriate documentation as/when they become available, and to contribute
24 towards the coordination of *focussed* experiments (see Sect. 3) that will test specific hypotheses
25 associated with model and climate sensitivity to the location, duration and magnitude of
26 freshwater fluxes.

27 **2.6 Topography, bathymetry, coastlines and rivers**

28 Changes in the ice sheets and their glacial eustatic and isostatic influence affected continental
29 topography and ocean bathymetry, which in turn shifted the coordinates of river mouths and
30 the coastal outline throughout the deglaciation. Hence time-varying topographic, bathymetric

1 and land-sea mask fields that match the chosen ice sheet from Sect. 2.4 (i.e. ICE-6G_C or
2 GLAC-1D) should be used; [these are provided within the ice sheet reconstruction datasets](#).

3 Topography should be updated at the same time as the model's ice sheet is updated; this is
4 mainly implicit to implementing the ice sheet reconstruction because the major orographic
5 changes through the deglaciation relate directly to ice sheet evolution. This said, due to glacial
6 isostatic adjustment components in the ice sheet reconstructions, there is evolution in
7 continental topography that is not directly the lowering/heightening of the ice surface, and it is
8 up to individuals whether they incorporate this or mask only the changes in ice sheet orography.

9 Ocean bathymetry will be provided, ~~but is an optional~~. [When deemed possible, this](#) boundary
10 condition ~~to vary~~[should be varied](#) through time. ~~Coastlines~~, [Where differences in the land-sea](#)
11 [mask require extra land to fill up coastal regions, or land to be cut away into ocean as sea level](#)
12 [rises \(see next paragraph on the other hand, coastlines\), the model must be changed accordingly,](#)
13 [because it is important to adequately represent the changing land-sea mask; for example, in](#)
14 [order to include overlying grounded ice.](#)

15 [Following on from this, coastlines](#) will need to be varied according to changes in global sea
16 level (and each model's horizontal grid resolution). It will be left to the discretion of participants
17 to decide how often to update either boundary condition, and when deciding on their frequency
18 it is recommended that groups consider the implications for opening/closing seaways and their
19 effect on ocean circulation and climate. Furthermore, the frequency need not be regular and
20 may instead focus on key 'events' in the marine [gateway] realm. However, whenever possible
21 and foreseeable, groups are encouraged to avoid making stepwise changes to model boundary
22 conditions that would interfere with signals of abrupt climate change; particularly those events
23 that the working group aims to focus on (e.g. Heinrich Event 1, the Bølling Warming, MWP1a,
24 the Younger Dryas etc.) unless the forcing (e.g. opening of a gateway) is assumed to be linked
25 with the event.

26 If groups wish, model river networks can be remapped to be consistent with this and updated
27 on the same timestep as the ice sheet reconstruction, either manually or by the model. However,
28 it is appreciated that the technical challenges associated with such a methodology would be
29 impractical for many. Therefore, following the recommendation of the PMIP3 LGM Working
30 Group (2010) [and Kageyama et al. \(in prep.\)](#), 'river pathways and basins should be at least
31 adjusted so that fresh water is conserved at the Earth's surface and care should be taken that

rivers reach the ocean' at every timestep that the bathymetry is adjusted; for example, when sea levels were lower, some river mouths may need to be displaced towards the [new] coastline to make sure they reach the ocean.

2.7 Vegetation, land surface and other forcings

In this section, recommendations are made for last deglaciation vegetation, land surface and aerosol (dust) parameters in the model.

There are three recommended options for setting up the Core ~~simulation's~~^{experiment's} vegetation and land surface parameters, they can either be: (i) computed using a dynamical vegetation model (e.g. coupled to the atmospheric component of the model); (ii) prescribed to match the CMIP5 preindustrial setup (Taylor et al., 2011a, 2011b) with fixed vegetation types and fixed plant physiology (including leaf area index); or (iii) prescribed to match the CMIP5 preindustrial setup (Taylor et al., 2011a, 2011b) with fixed vegetation types and interactive plant physiology if running with an enabled carbon cycle. If prescribing vegetation and land surface, i.e. using option (ii) and (iii), groups should be aware that coastal land will be emerged compared to preindustrial because of the increased terrestrial ice volume and associated lower eustatic sea level (with the maximum during the early stages of the Core). Therefore, vegetation/land surface will need to be interpolated onto the emerged land from preindustrial grid cells, for example using nearest neighbour methods.

For models with prognostic aerosols, the parameters for dust [forcing] can be computed dynamically. Alternatively, it is recommended that Core simulations fix the associated parameters according to the CMIP5 preindustrial simulation (Taylor et al., 2011a, 2011b), ~~with no temporal variation.~~ with no temporal variation. Examining the influence of different transient aerosol scenarios (for those models that do not include prognostic dust, for example) could constitute a further suite of sensitivity simulations for comparison with the Core

~~It has already been described that for the LGM (i.e. the very start of the Core simulation), groups are recommended to adjust the global freshwater budget by +1 psu to account for the increased [terrestrial] ice volume (Sect. 2.1.1). If salinity is reset at any subsequent point (e.g. to correct for model drifts or to account for ice volume changes), this must be documented.~~

1 There is no last deglaciation protocol for setting up other forcings, transient or fixed in time.
2 For all simulations, groups are required to fully document their methods, including experiment
3 design and especially when different or with additional components to the setup described here.

4 **3 Coordinating further simulations**

5 As already ~~alluded to~~discussed, we are faced with the challenge of designing an experiment that
6 is suitable to be run with a wide range of models, from the more computationally efficient class
7 of intermediate complexity models, to state-of-the-art Earth System Models. One particular
8 difficulty is enabling the most complex and highest resolution climate models to participate in
9 this 12 thousand year long experiment when for some, even the integration to reach the LGM
10 spinup state demands a huge amount of computational resource. There is no easy solution and
11 our approach will be to augment the Core ~~simulations~~simulations with shorter *focussed*
12 simulations that target specific questions, mechanisms and time periods. Whilst the most
13 computationally expensive models (e.g. the latest generation of Earth System Models) may not
14 ~~feasibly~~ be able to participate in the Core, they will be included in the shorter subset of *focussed*
15 simulations. Similarly, alternative full-deglaciation simulations can be coordinated for the less
16 computationally expensive models in the working group (e.g. low resolution General
17 Circulation Models, and Earth System Models of Intermediate Complexity).

18 One line of investigation relating to meltwater inputs from ice sheets and icebergs is to carry
19 out a suite of sensitivity simulations examining different injection sites. These simulations
20 would help to address some of the uncertainty ~~that led to the exclusion of~~in freshwater ~~fluxes~~
21 ~~from the Core~~flux scenarios. For example, geochemical evidence suggests that smaller and
22 more localised discharges of freshwater than have traditionally been considered in climate
23 models may have an important influence on ocean circulation (e.g. Hall et al., 2006), implying
24 that precise freshwater fluxes are needed in the models to examine their effect. Certainly, others
25 have shown that the location of injection is a controlling factor on the impact of freshwater
26 delivery to the ocean, not just laterally (e.g. Condrón and Winsor, 2012; Smith and Gregory,
27 2009), but also in terms of depth (e.g. Roche et al., 2007).

28 A set of coordinated simulations exploring a range of uncertainty in the freshwater forcing
29 (location, depth, duration, magnitude, and physical characteristics such as temperature and
30 density) would be well suited for the *focussed* experiments, thus building on the ~~meltwater-free~~

~~Core.~~ Core simulations, which may themselves indicate interesting avenues for investigation; partly the purpose of a flexible meltwater approach.

However, freshwater is not the only issue and other *focussed* experiments could include the influence of ~~timing in~~ greenhouse gas ~~records, record~~, differences in ice sheet reconstructions (e.g. the PMIP3 merged ice sheet from Abe-Ouchi et al., 2015; ICE-6G_C; GLAC-1D) or simulations with [coupled] ice-sheet models, the relative importance of different forcings (e.g. insolation vs. trace gases vs. ice sheet evolution), sensitivity to dust-forcing scenarios, the influence of changes in tidal energy dissipation (Schmittner et al., 2015), event-specific hypothesis testing, and shorter-term variability within the climate system.

Based on on-going discussions, it is likely that the first ~~sets~~ of *focussed* simulations will be:

- Sensitivity and hypothesis-driven, investigating simulations that compare results from uniformly distributed meltwater fluxes to results from river-routed meltwater fluxes to examine the impact of the regional specificity of freshwater forcing upon climate system evolution.
- Sensitivity simulations that are free from ice meltwater fluxes to provide information on what climate evolution was caused by processes other than freshwater fluxes to the ocean.
- A hypothesis-driven investigation of the possible mechanisms for preconditioning the glacial ocean for the relatively cool Heinrich Stadial 1 and ensuing catastrophic iceberg discharge (Barker et al., 2015).
- Sensitivity experiments examining the role of trace gas forcing resolution on climate evolution; for example, smoothing the record provided by Bereiter et al. (2015).

We have described the plans for *focussed* simulations to highlight the depth of the working group's aims and to properly contextualise the Core ~~simulations~~ simulations, but the purpose of this manuscript is to outline the model setup for the Core ~~simulation. The~~ experiment. The design for subsequent *focussed* simulations will be described at a later date on the PMIP Last Deglaciation Working Group Wiki (<https://wiki.lsce.ipsl.fr/pmip3/doku.php/pmip3:wg:degla:index>) and we welcome contributions to the discussion of what further simulations to coordinate there.

4 Working group phases

The experiment will be split into three phases that are designed to run seamlessly into each other (Fig. 1a). Phase one begins at the LGM (21 ka) and will finish at the abrupt Bølling Warming event, which is where Phase 2 picks up, encompassing the Bølling Warming. Phase 3 begins at the start of the Younger Dryas cooling and is currently planned to continue through to the end of the Core ~~simulation~~experiment at 9 ka.

Perhaps most importantly, this affords near-future milestones for managing the ultimate completion of the long full deglacial simulation across all participant groups. It will provide a timetabled framework for beginning and continuing the longer simulations; for scheduling shorter, event- or challenge-specific transient simulations by more computationally expensive models (see discussion in Sect. 3); and for the analysis and publication of results as the milestones are reached. Another motivation is to ensure that the experiment design for later periods of the last deglaciation is updated according to knowledge gained from simulations of the preceding time period; for example, changes in ocean and climate states, which have previously been shown to have a strong influence on climate trajectories (e.g. Kageyama et al., 2010; Timm and Timmermann, 2007). This is particularly important for setting up shorter, event-specific focussed simulations, but it is not planned to be explicitly used to influence the Core. Splitting the period into phases also provides the opportunity to update model boundary conditions and climate forcing data with cutting edge palaeoclimate reconstructions, as they emerge during the lifespan of the multi-model experiment. However, care will be taken to ensure that these are physically consistent between phases—, and these updates will not compromise the Core simulations described in this manuscript. This is so as not to disadvantage more computationally efficient models that may have already completed simulating the full 21-9 ka (or beyond) period. Instead, the information will be incorporated into focussed versions of the last deglaciation simulations; possibly spun-off sub-periods that do not have to start again at the LGM.

Each phase will encompass at least one distinguishable climate event; Heinrich Stadial 1 and Heinrich Event 1 in Phase 1 following on from the LGM; MWP1a, the Bølling Warming and the Antarctic Cold Reversal in Phase 2; and the Younger Dryas cooling in Phase 3 (Fig. 1b). As outlined in Sect. 3, simulations of these shorter events can be coordinated in the *focussed* simulations. This is to engage the higher complexity/resolution models, which are unable to run

longer simulations, but can use the wider framework of the working group to provide valuable knowledge on rapid climate changes known to have taken place in the last 21 ka.

5 Summary

The last deglaciation presents a host of exciting opportunities to study the Earth System and in particular, to try to understand a range of abrupt climate changes that occurred over just a few years to centuries within the context of more gradual trends. Numerical climate models provide useful tools to investigate the mechanisms that underpin the events of this well-studied time period, especially now that technological and scientific advances make it possible to run multi-millennium simulations with some of the most complex models. Several recent modelling studies have begun this task, but many questions and untested hypotheses remain. Therefore, under the auspices of the Paleoclimate Modelling Intercomparison Project (PMIP), we have set up an initiative to coordinate efforts to run transient simulations of the last deglaciation, and to facilitate the dissemination of expertise between modellers and those engaged with reconstructing the climate of the last 21 thousand years.

The first step has been to design a ~~single~~, Core ~~simulation~~experiment suitable for a range of PMIP models; from relatively fast and coarse resolution Earth System Models of Intermediate Complexity, to new generations of the more complex and higher resolution General Circulation and Earth System Models. The setup for this Core ~~simulation~~experiment, is based on an approach that tries to combine a traditional Model Intercomparison Project method of strictly prescribing boundary conditions across all models, and the philosophy of utilising the breadth of participants to address outstanding uncertainty in the climate forcings, model structure and palaeoclimate reconstructions. Accordingly, we have made recommendations for the initialisation conditions for the simulation and have stated our minimum requirements for the transient experiment design, as summarised in Table 1 and 2, respectively.

However, there are some uncertainties that the Core is not designed to deal with directly or exhaustively; two examples discussed in this manuscript being the effect of trace gas record resolution and the influence of ice melt on the oceans and climate, ~~and the effect of timing in the trace gas records respectively~~. We know that the Core ~~simulations~~simulations will not tackle all of our questions, and ~~is~~are likely to give rise to others. Therefore, additional *focussed* simulations will also be coordinated on an ad-hoc basis by the working group. Many of these

will build on and be centred around the Core; often taking shorter snapshots in time, thus including the most computationally expensive models in the experiment, or presenting twelve-thousand year alternatives to the Core for faster models to contribute. Not all simulations will be suitable for all models, but the aim is that taken as a whole, the experiment can utilise the wide range of PMIP model strengths and hence minimise individual weaknesses.

Essentially, the Core [simulationexperiment](#) has been designed to be inclusive, taking into account the best compromise between uncertainties in the geological data and model limitations. The hypothesis-driven *focussed* experiments will go further than the Core to target the questions that remain. It is hoped that this exciting initiative will improve our individual efforts, providing new opportunities to drive the science forwards towards understanding this fascinating time period, specific mechanisms of rapid climate warming, cooling and sea level change, and Earth's climate system more broadly.

Author Contributions

RFI and LJG lead the PMIP Last Deglaciation Working Group, for which AB, MK, DMR and PJV act as the advisory group. RFI, LJG, MK, DMR, PJV and AB collaboratively designed the working group's aims, structure, Core [simulationexperiment](#) and additional experiments in consultation with the wider community. RD, WRP and LT provided the ice sheet reconstructions, plus associated boundary conditions. RFI and LJG collated these and all other boundary condition data for the simulations. RFI and LJG wrote the manuscript and produced the figures with contributions from all authors.

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Tables

Table 1. Summary of recommended model boundary conditions to spin up the last deglaciation Core [simulationexperiment](#) (pre 21 ka); see text for details. Participants are not required to follow the recommendation for these boundary conditions, but must document the method used, including information on the simulation's state of spinup at the point when the Core is started. Data are available from PMIP Last Deglaciation Working Group Wiki: <https://wiki.lsce.ipsl.fr/pmip3/doku.php/pmip3:wg:degla:index>. Boundary condition group headings are in bold.

Spinup type	Boundary condition	Description
Last Glacial Maximum (LGM; 21 ka)	Insolation	
	Solar constant	Preindustrial (e.g. 1365 W m ⁻²)
	Eccentricity	0.018994
	Obliquity	22.949°
	Perihelion–180°	114.42°
	Vernal equinox	Noon, 21 st March
	Trace gases	
	Carbon dioxide (CO ₂)	488 190 ppm
	Methane (CH ₄)	375 ppb
	Nitrous oxide (N ₂ O)	200 ppb
Transient orbit and trace gases (26-21 ka)	Chlorofluorocarbon (CFC)	0
	Ozone (O ₃)	Preindustrial (e.g. 10 DU)
	Ice sheets, orography and coastlines	21 ka data from either: - ICE-6G_C (references in text) - GLAC-1D (references in text)
	Bathymetry	Keep consistent with the coastlines, using either: - Data associated with the ice sheet - Preindustrial bathymetry
	Global ocean salinity	+ 1 psu, relative to preindustrial
	Orbital parameters	All orbital parameters should be transient, as per Berger (1978) 26-21 ka
	Trace gases	Adjusted to the AICC2012 (Veres et al., 2013)
	Carbon dioxide (CO ₂)	Transient, as per LüthiBereiter et al. (20082015)
	Methane (CH ₄)	Transient, as per Louergue et al. (2008)
	Nitrous oxide (N ₂ O)	Transient, as per Schilt et al. (2010)
	All others	As per LGM (21 ka) spinup type.

Table 2. Summary of required model boundary conditions for the last deglaciation Core simulation experiment 21-9 ka; optional boundary conditions are labelled as such. Data are available from PMIP Last Deglaciation Working Group Wiki: <https://wiki.lsce.ipsl.fr/pmip3/doku.php/pmip3:wg:degla:index>. See text for details. Boundary condition group headings are in bold.

Boundary condition	Description
Initial conditions (pre 21 ka)	Recommended (optional) to use either: <ul style="list-style-type: none"> - Last Glacial Maximum (LGM; 21 ka) equilibrium simulation, including +1 psu global ocean salinity - Transient orbit and trace gases (26-21 ka) and all other boundary conditions fixed as per equilibrium LGM See Table 1 for details. The method must be documented, including information on the state of spinup
Insolation	
Solar constant	Preindustrial (e.g. 1365 W m ⁻²)
Orbital parameters	Transient, as per Berger (1978)
Trace gases	Adjusted to the AICC2012 age model (Veres et al., 2013):
Carbon dioxide (CO ₂)	Transient, as per Lüthi+Bereiter et al. (2015)
Methane (CH ₄)	Transient, as per Louergue et al. (2008)
Nitrous oxide (N ₂ O)	Transient, as per Schilt et al. (2010)
Chlorofluorocarbon (CFC)	0
Ozone (O ₃)	Preindustrial (e.g. 10 DU)
Ice sheet	Transient, with a choice of either : <ul style="list-style-type: none"> - ICE-6G_C reconstruction (references in text) - GLAC-1D reconstruction (references in text) How often to update the ice sheet is optional
Orography and coastlines	Transient. To be consistent with the choice of ice sheet. Orography is updated on the same timestep as the ice sheet. It is optional how often the land-sea mask is updated, but ensure consistency with the ice sheet reconstruction is maintained
Bathymetry	Keep consistent with the coastlines and otherwise use either: <ul style="list-style-type: none"> - Transient data associated with the chosen ice sheet; it is optional how often the bathymetry is updated- - Preindustrial bathymetry
River routing	Ensure that rivers reach the coastline It is recommended (optional) to use one of the following: <ul style="list-style-type: none"> - Preindustrial configuration for the model - Transient routing provided with the GLAC-1D ice sheet reconstructions - Manual/model calculation of river network to match topography
Freshwater fluxes	No land ice or iceberg meltwater fluxes to the ocean At participant discretion. Three options are: melt-uniform, melt-routed and no-melt (see text). It is recommended (optional) to run at least one Core simulation with a scenario consistent with the chosen ice sheet reconstruction to conserve salinity (e.g. as provided). See text for full details (Sect. 2.5)
Other (optional)	
Vegetation and land cover	Prescribed preindustrial cover or dynamic vegetation model
Aerosols (dust)	Prescribed preindustrial distribution or prognostic aerosols

Figures

Figure 1. The last deglaciation; forcings and events. **(a)** The three phases of the Core simulation experiment (Sect. 4). **(b)** Climate events/periods discussed in the text; Last Glacial Maximum (LGM; 23-19 ka as according to the EPILOG definition; Mix et al., 2001), Heinrich Stadial 1 (HS1), Heinrich Event 1 (H1), Bølling Warming (BW) and Meltwater Pulse 1a (MWP1a), Antarctic Cold Reversal (ACR) and the Younger Dryas cooling (YD). **(c)** June insolation at 60° N and December insolation at 60° S (Berger, 1978). **(d)** Atmospheric carbon dioxide concentration (recent composite of EPICA Dome C, Vostok, Taylor Dome, Siple Dome and West Antarctic Ice Sheet Divide records, Antarctica; Bereiter et al., 2015); black dashed line shows preindustrial concentration. **(e)** Atmospheric methane concentration (EPICA Dome C, Antarctica; Loulergue et al., 2008); green dashed line shows preindustrial concentration. **(f)** Atmospheric nitrous oxide concentration (Talos Dome, Antarctica; Schilt et al., 2010); brown dashed line shows preindustrial concentration. **(g)** Volume of the ice sheets according to the ICE-6G_C reconstruction (solid lines; Argus et al., 2014; Peltier et al., 2015) and the GLAC-1D reconstruction (dashed lines; Briggs et al., 2014; Tarasov et al., 2012; Tarasov and Peltier, 2002). Associated meltwater scenarios *melt-uniform* and *melt-routed* (see Sect. 2.5) are consistent with these; all ice mass loss shown is supplied as freshwater to the ocean. **(h)** Greenland temperature reconstruction with $\pm 1 \sigma$ shaded (averaged GISP2, NEEM and NGRIP records; Buizert et al., 2014). **(i)** Antarctic δD (EPICA Dome C; Jouzel et al., 2007). **(d)-(f)** and **(h)-(i)** are given on the AICC2012 timescale (Veres et al., 2013).

Figure 2. Northern Hemisphere ice sheet elevation at 21, 18, 15, 12 and 9 ka; **(a)** ICE-6G_C reconstruction at 10 arcminute horizontal resolution, elevation is plotted where the fractional ice mask is more than 0.5 (Peltier et al., 2015); **(b)** GLAC-1D reconstruction at 1° (*longitude*) \times 0.5° (*latitude*) horizontal resolution, elevation is plotted where the *binaryfractional* ice mask is *onemore than 0.5* (Briggs et al., 2014; Tarasov et al., 2012; Tarasov and Peltier, 2002; Tarasov et al., 2012; this study).

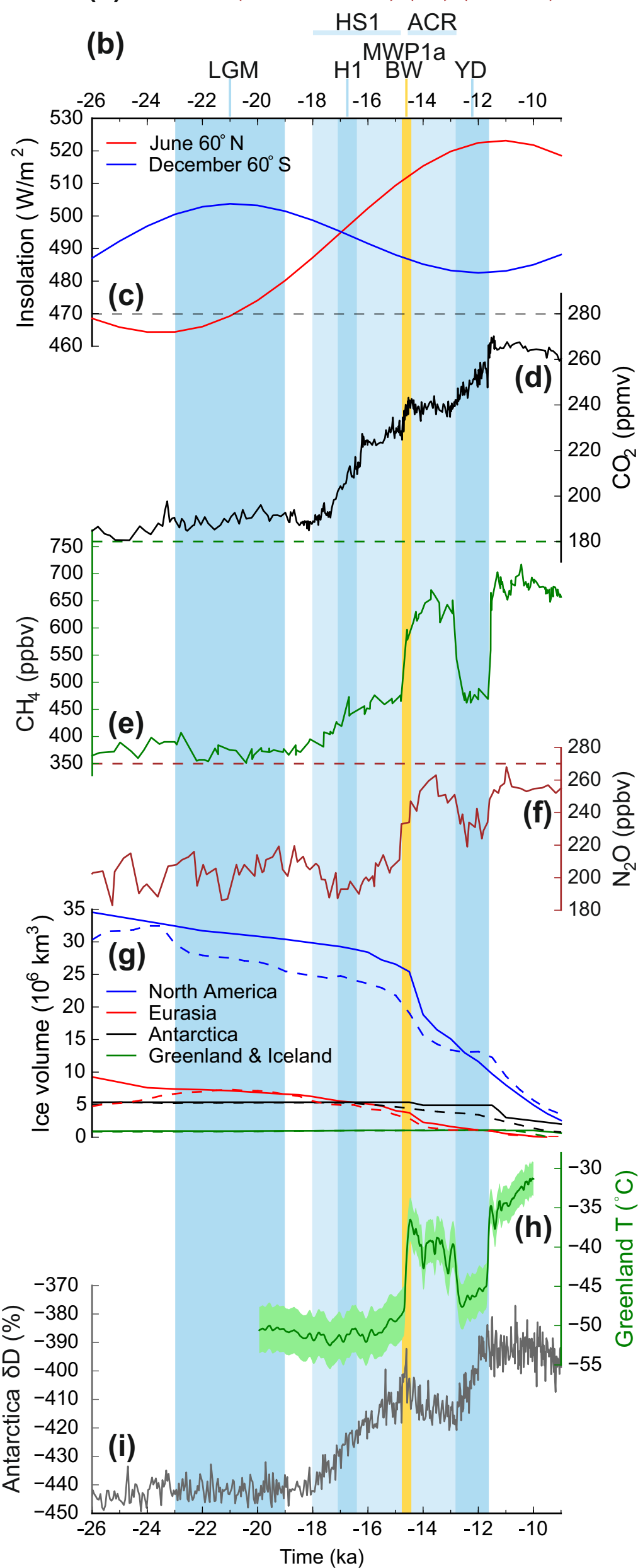
Figure 3. Atmospheric trace gases through the last deglaciation from Antarctic ice cores. **(a)** CarbonCore experiment carbon dioxide according to a recent composite record from EPICA Dome C (EDC), West Antarctic Ice Sheet Divide (WDC), Vostok and Taylor Dome and Siple Dome (thick black line; Bereiter et al., 2015), adjusted to, which was produced on the AICC2012 chronology (Veres et al., 2013). Also shown for comparison is an older composite

[record from EDC, Vostok and Taylor Dome](#) (thin blue line; Lüthi et al., 2008, adjusted to the AICC2012 chronology). ~~The, as well as the~~ original EDC CO₂ record (green line; Monnin et al., 2004) and ~~morethe~~ recent, higher resolution ~~West Antarctic Ice Sheet Divide (WDC)~~ CO₂ record (dark red line; Marcott et al., 2014); [which were incorporated into the newer composite by Bereiter et al. \(2015\)](#) ~~are shown for comparison.~~ (b) Methane according to the EPICA Dome C (EDC) record (Loulergue et al., 2008), shown both on the original EDC1 chronology (green line; Spahni et al., 2005) and adjusted to the more recent AICC2012 chronology [for the Core experiment](#) (thick black line; Veres et al., 2013). (c) Nitrous oxide according to the Talos Dome (TALDICE) record (Schilt et al., 2010), adjusted to the AICC2012 chronology [for the Core experiment](#) (thick black line; Veres et al., 2013). For comparison, the earlier EPICA Dome C (EDC) record on the EDC1 chronology is also shown (green line; Spahni et al., 2005). The nearest measured N₂O concentration to 21 ka is from 21.089 ka; hence the small offset between the slightly earlier concentration (187 ppb) used for the Core and the interpolated value plotted at 21 ka. For (a)-(c) 21 ka concentrations according to the AICC2012 age model (red dots) are shown in contrast to previous PMIP3 LGM concentrations (blue dots; PMIP LGM Working Group, 2010). If using an equilibrium-type spinup for the start of ~~thea~~ transient Core simulation at 21 ka (Sect. 2.1.1), use ~~488~~[190](#) ppm CO₂, 375 ppb CH₄ and 200 ppb N₂O.

Figure 4. Southern Hemisphere ice sheet elevation at 21, 12 and 9 ka; (a) ICE-6G_C reconstruction at 10 arcminute horizontal resolution, ice elevation is plotted where the fractional ice mask is more than 0.5 (Argus et al., 2014; Peltier et al., 2015); (b) GLAC-1D reconstruction at 1° [\(longitude\) × 0.5° \(latitude\)](#) horizontal resolution, [ice](#) elevation is plotted where ~~the binaryfractional~~ ice mask is ~~4~~[more than 0.5](#) (Briggs et al., 2014; Tarasov et al., 2012; Tarasov and Peltier, 2002).

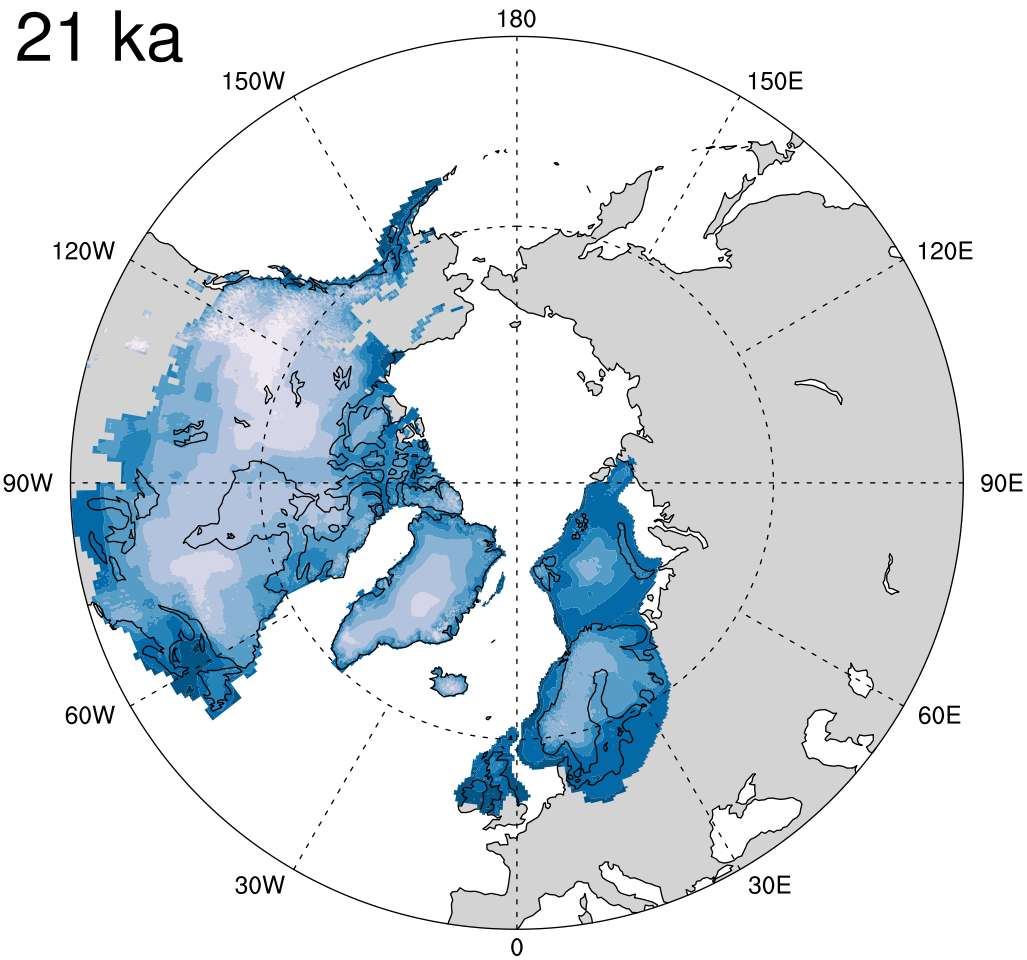
(a) Phases ← **1** → **2** → **3** →

(b)

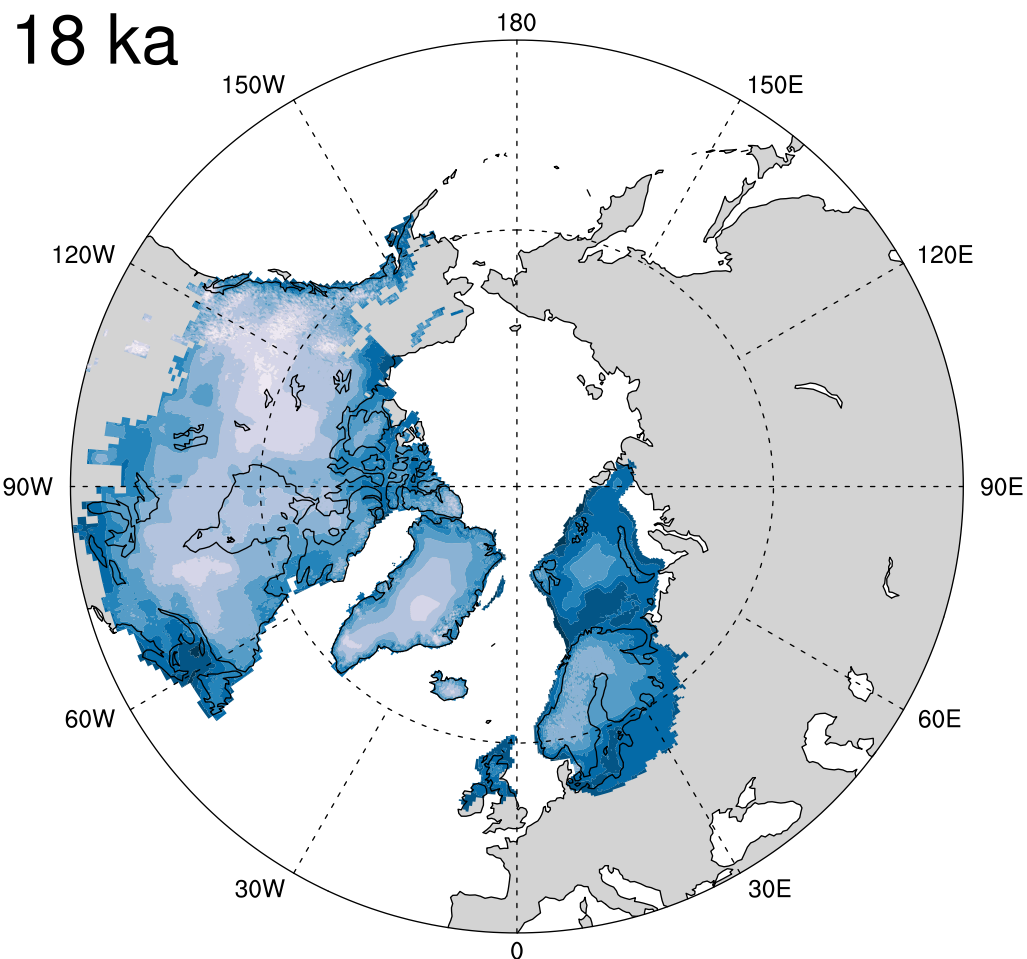


(a) ICE-6G_C

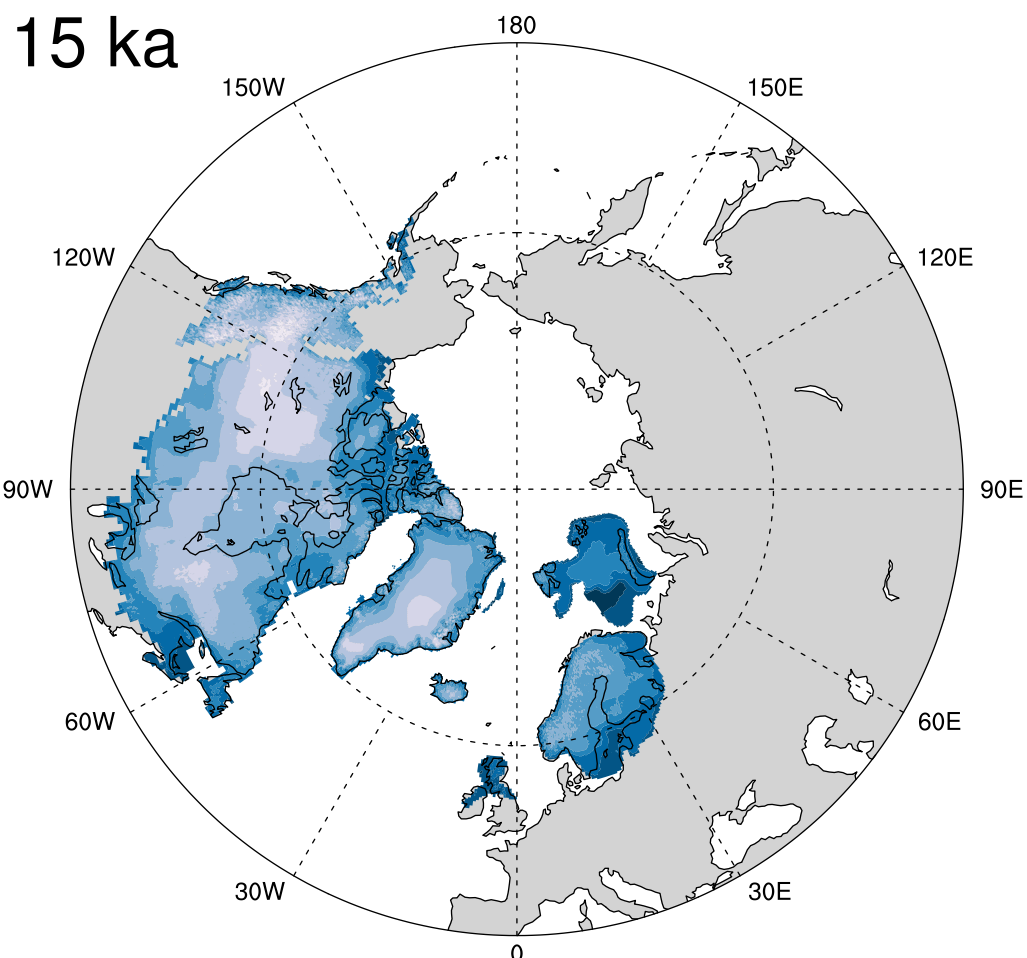
21 ka



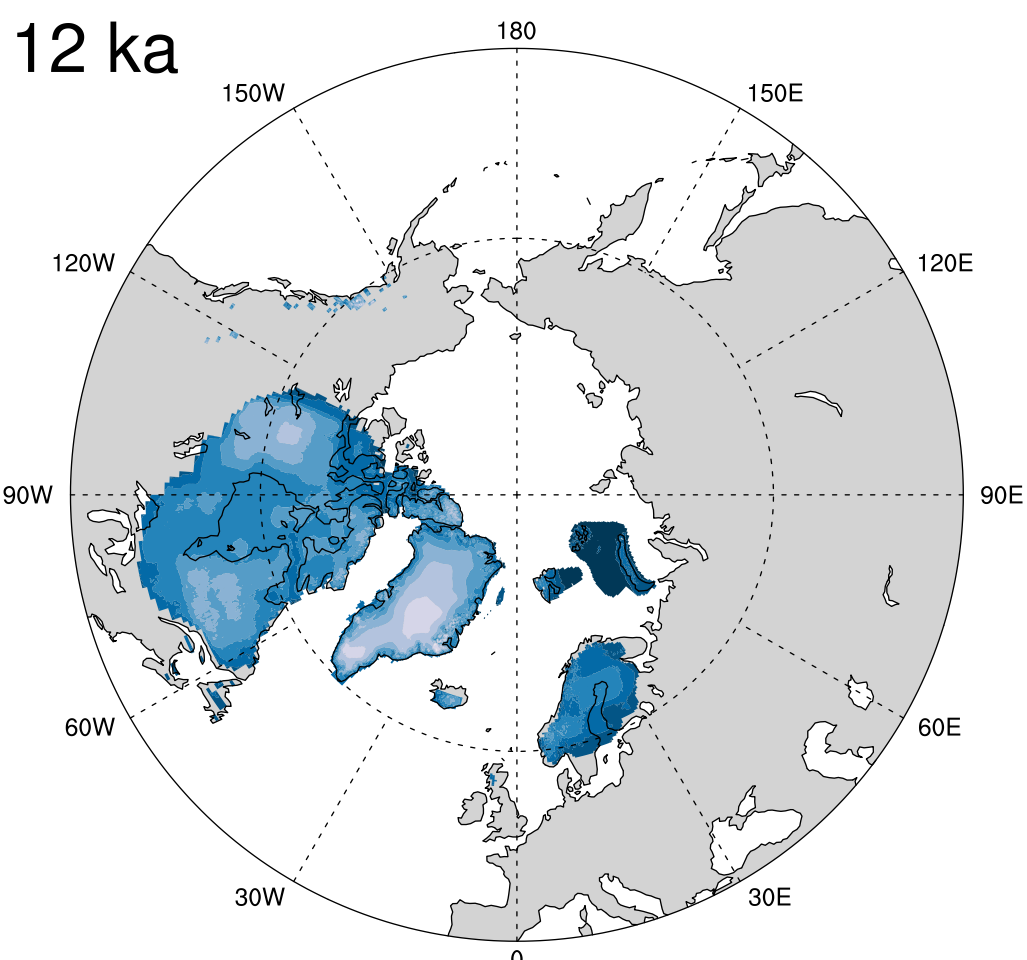
18 ka



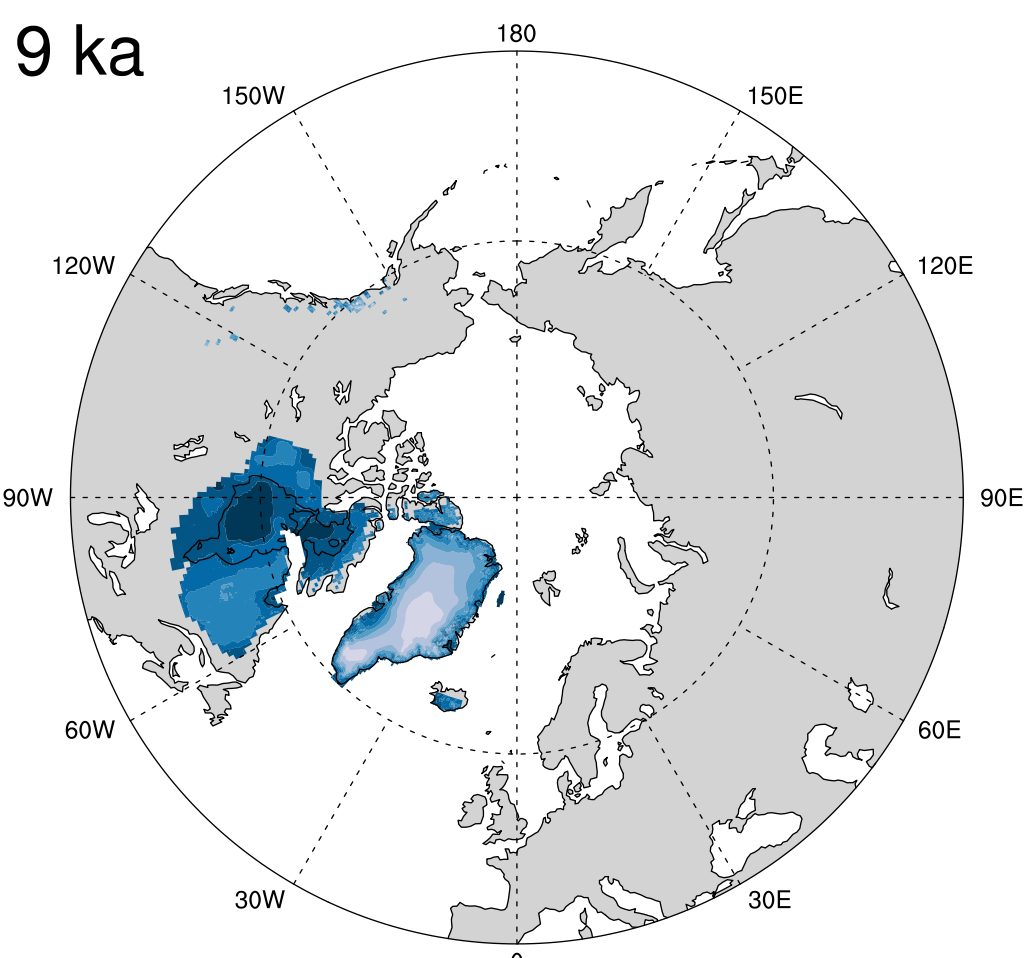
15 ka



12 ka

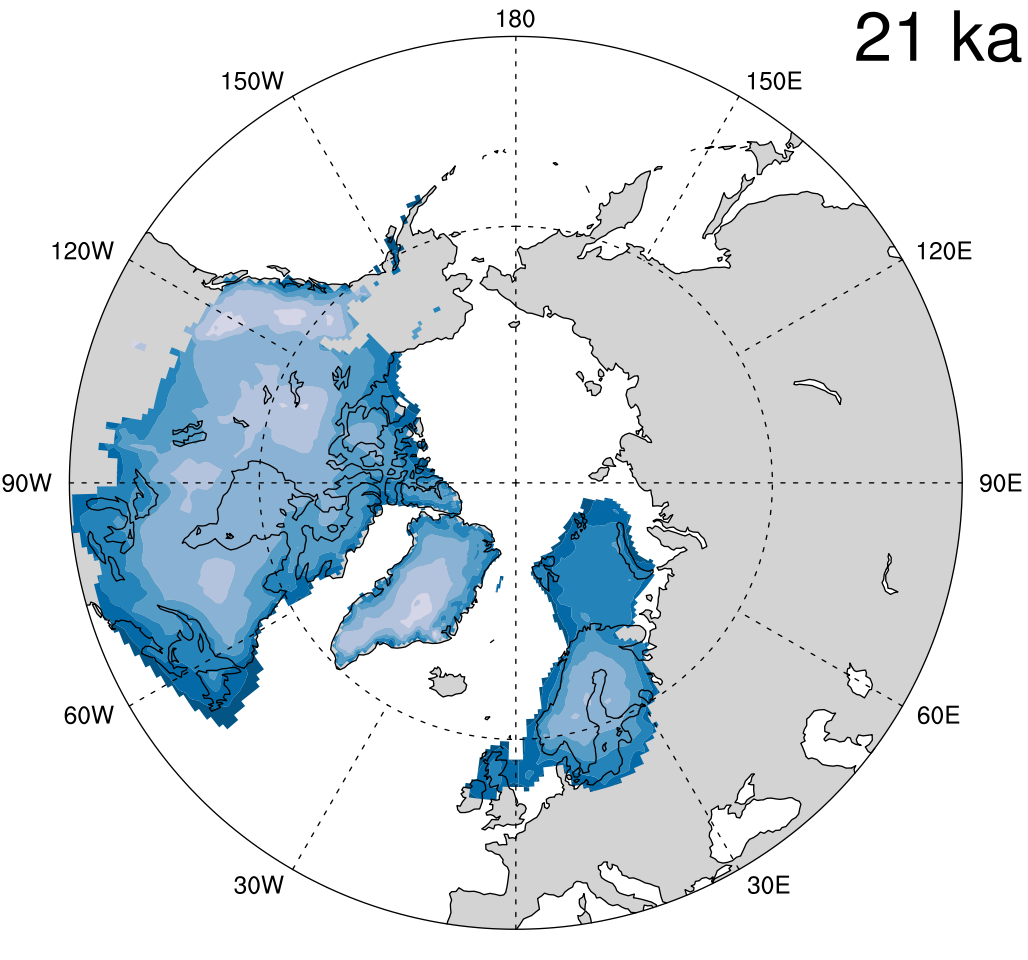


9 ka

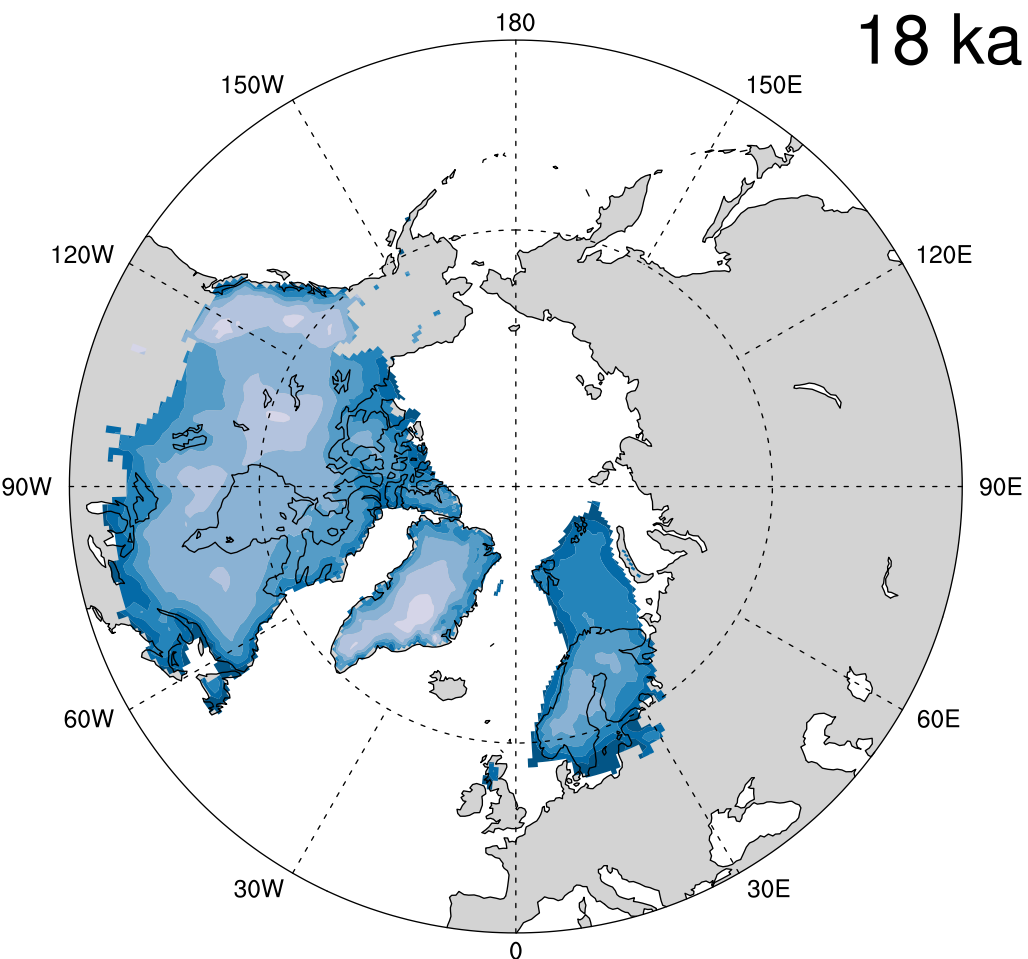


(b) GLAC-1D

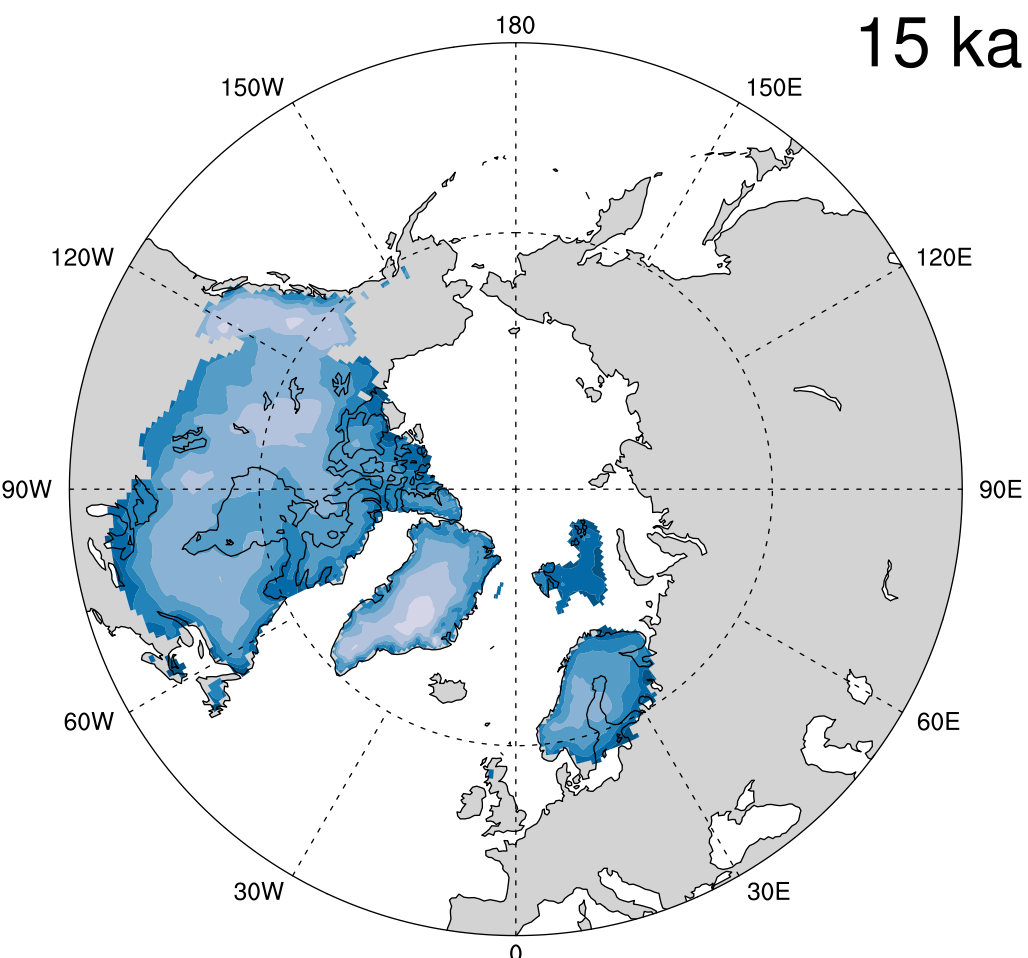
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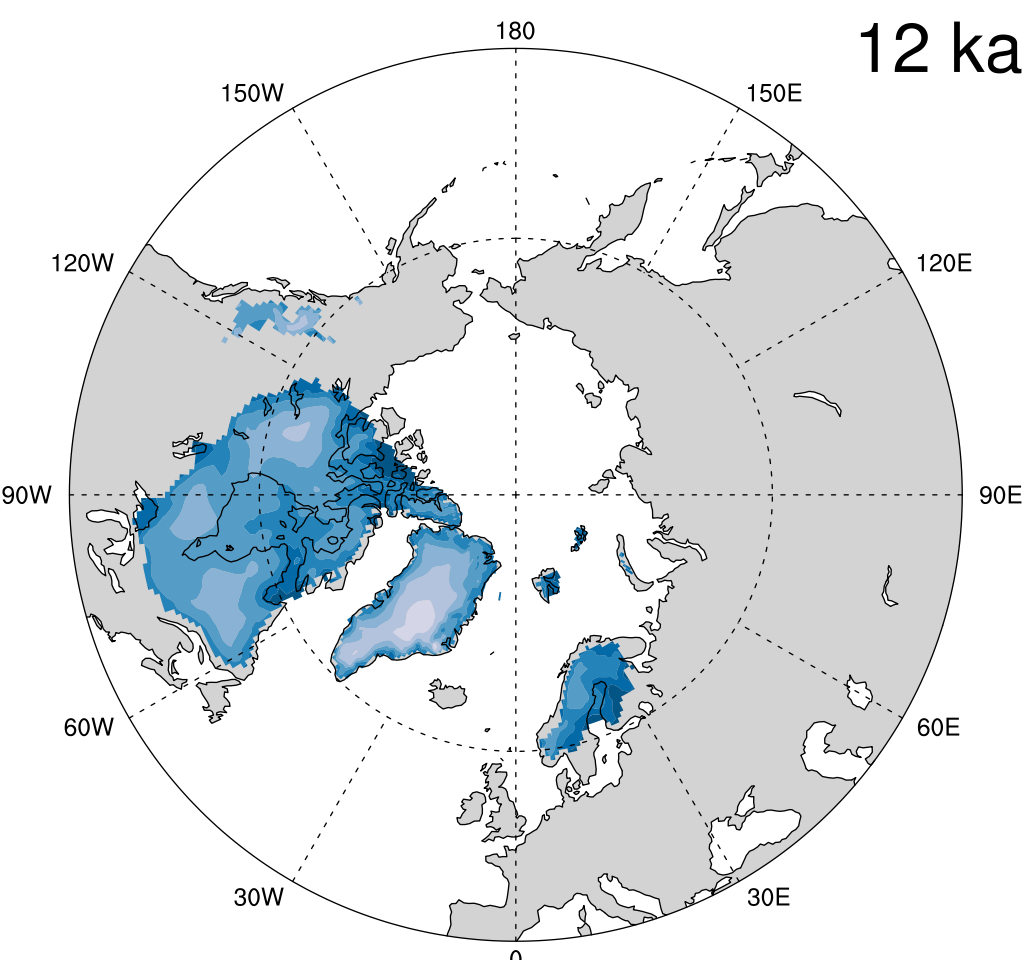
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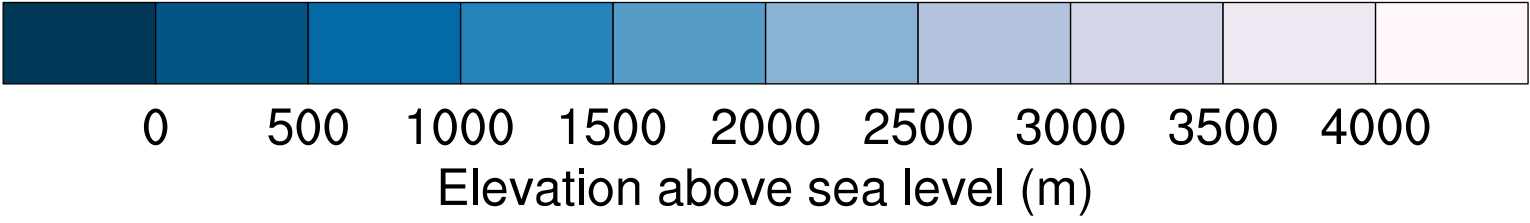
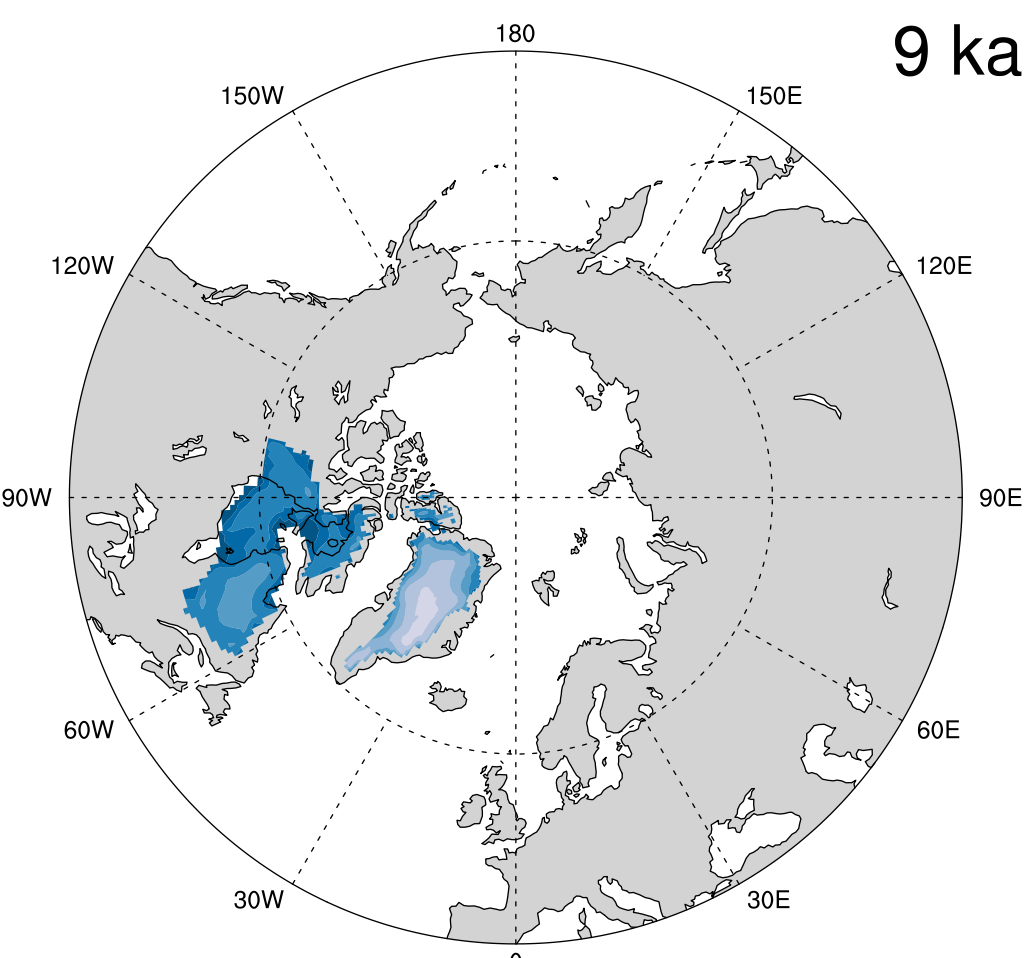
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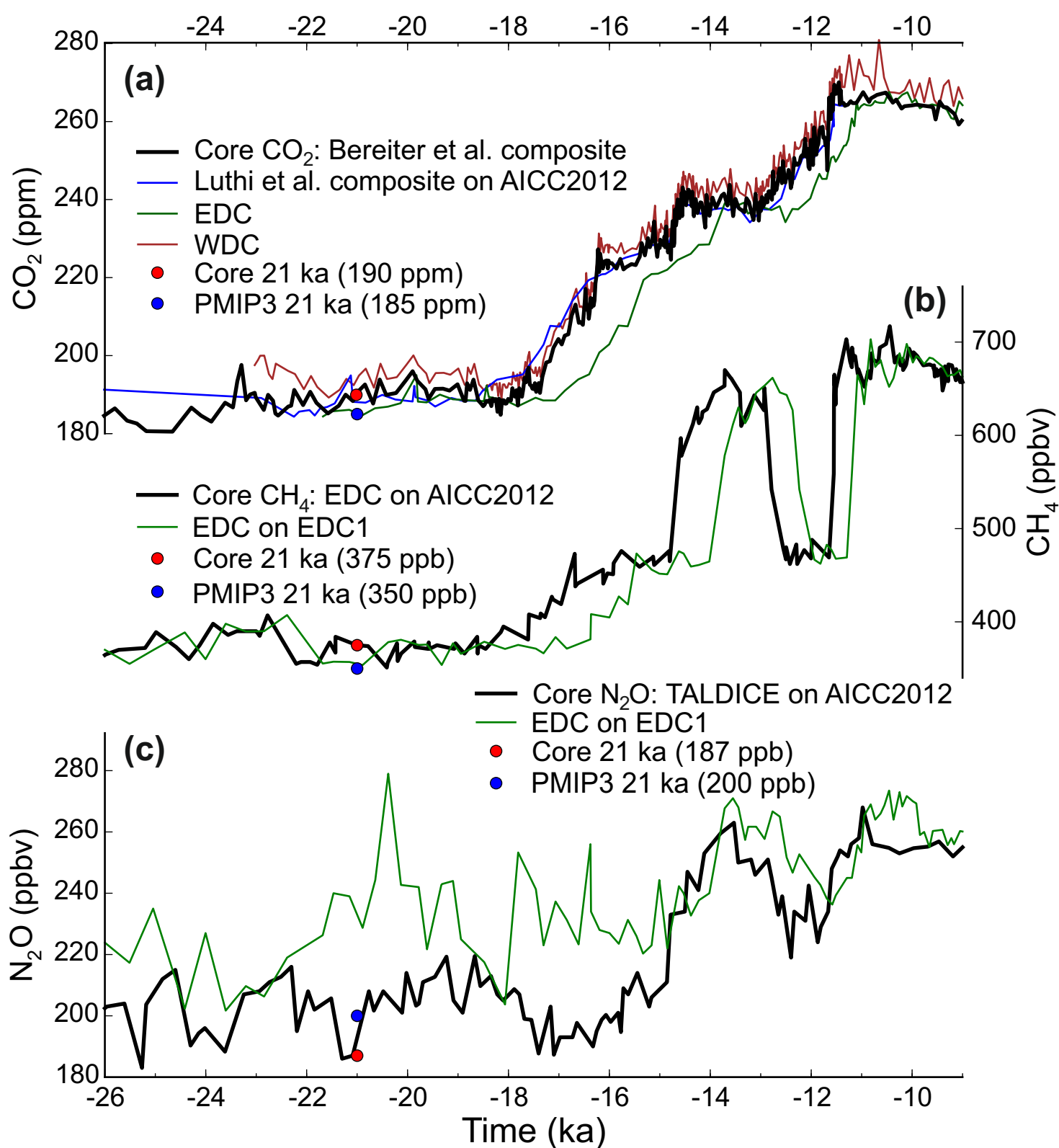


12 ka



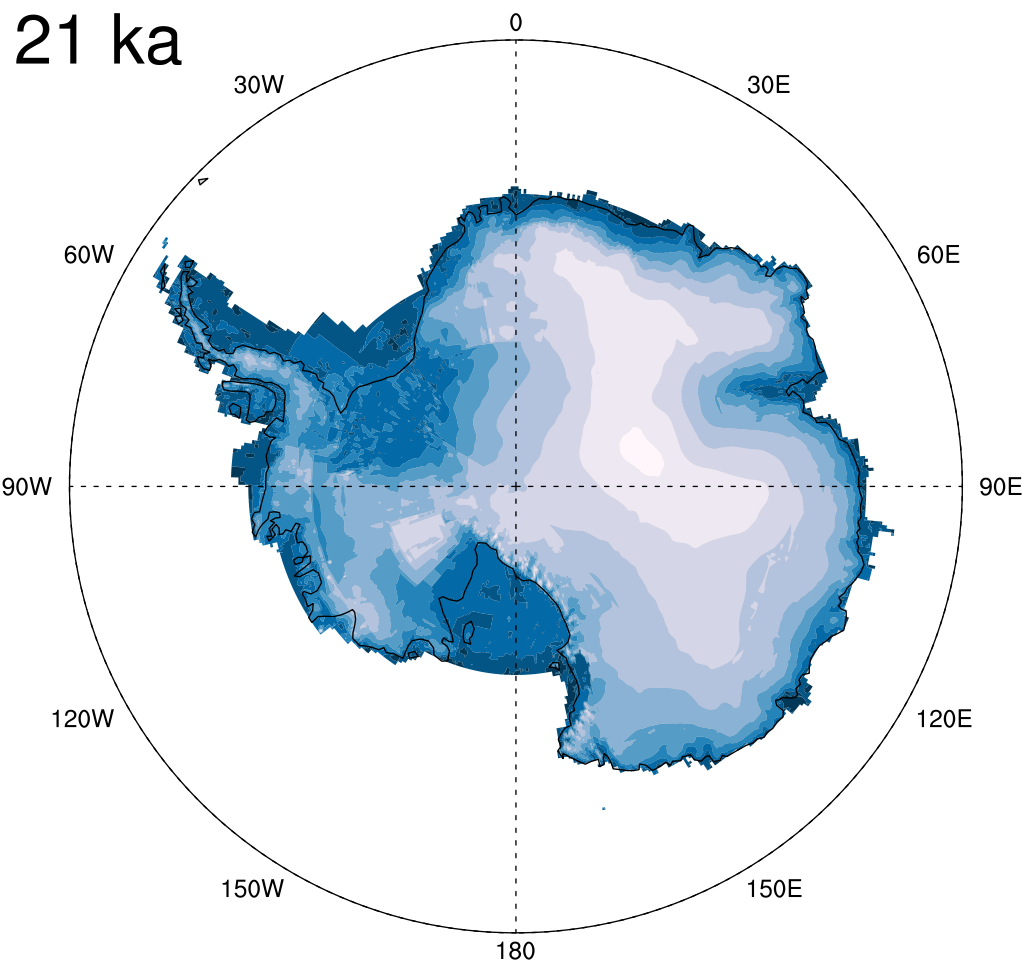
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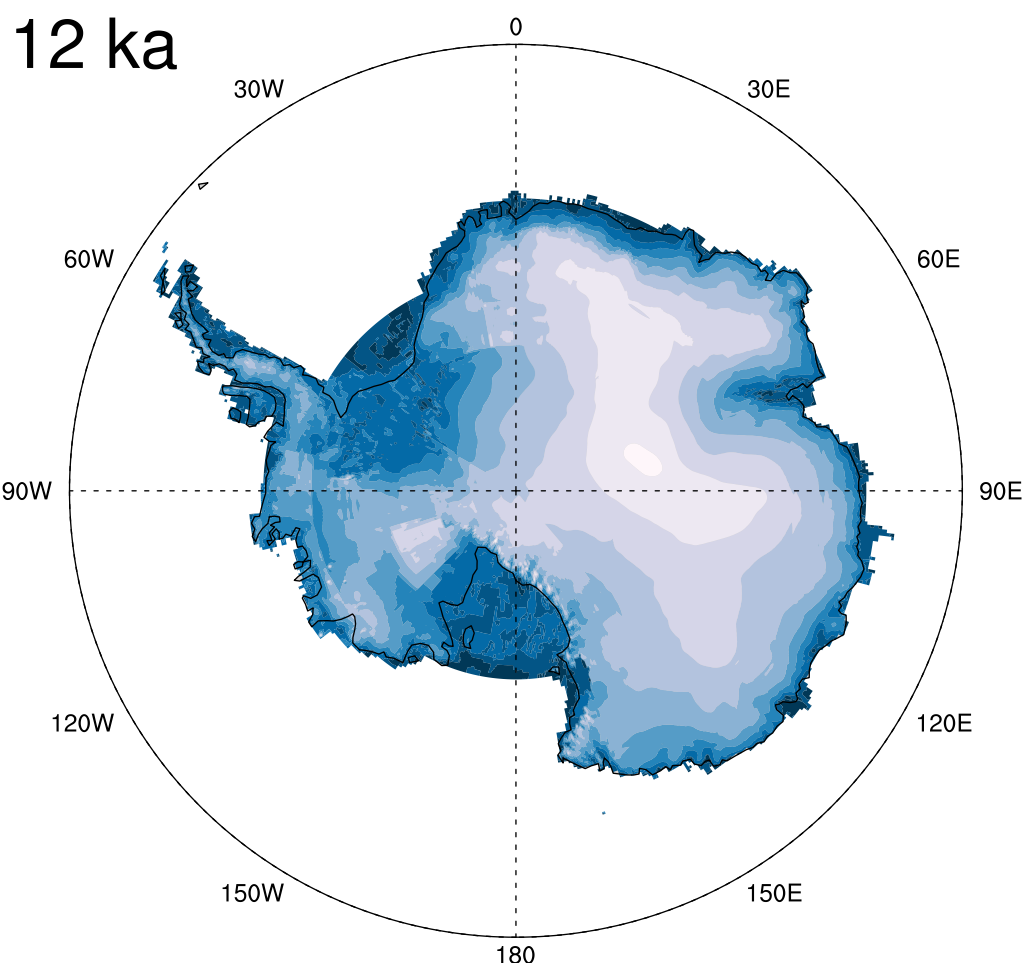


(a) ICE-6G_C

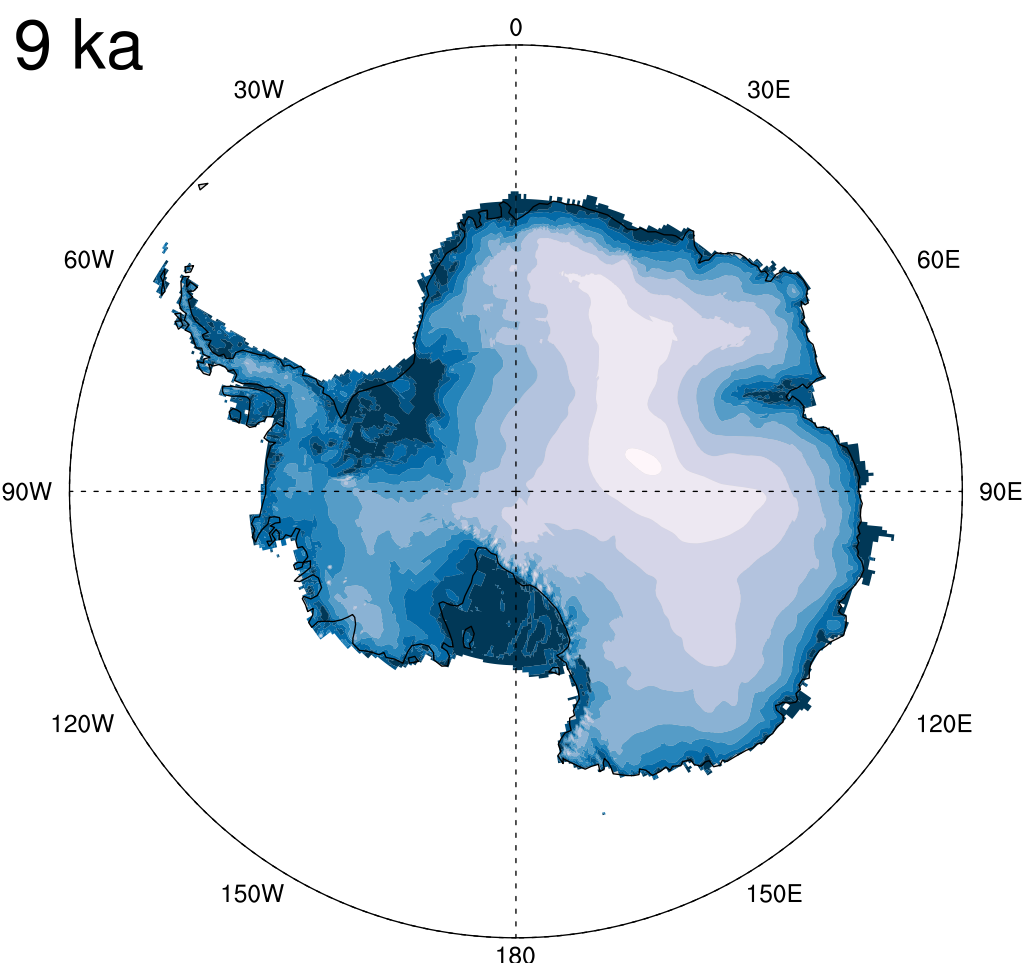
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12 ka

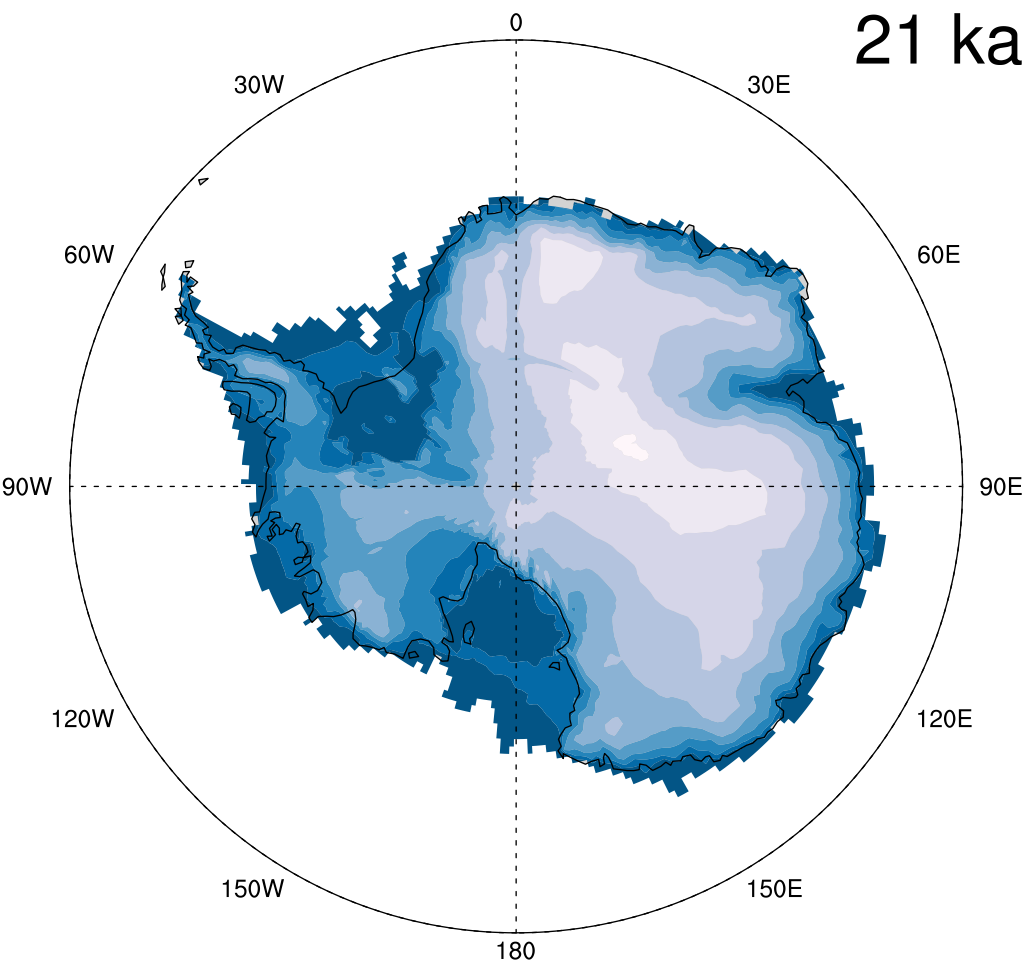


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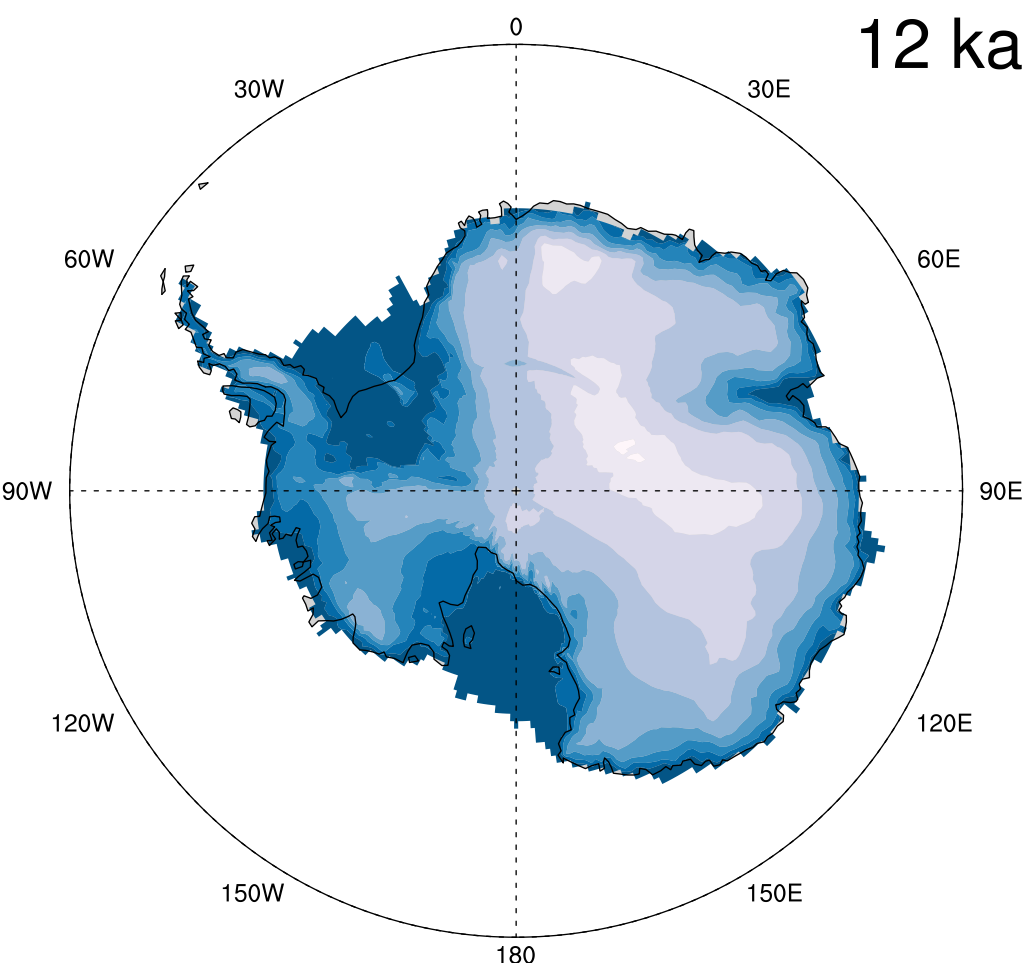


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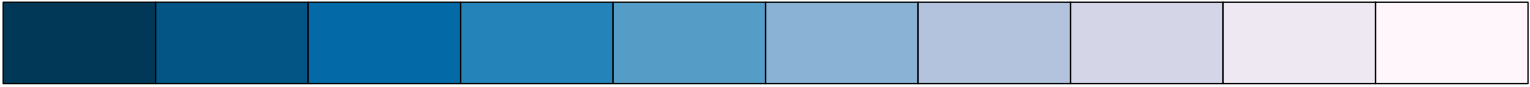
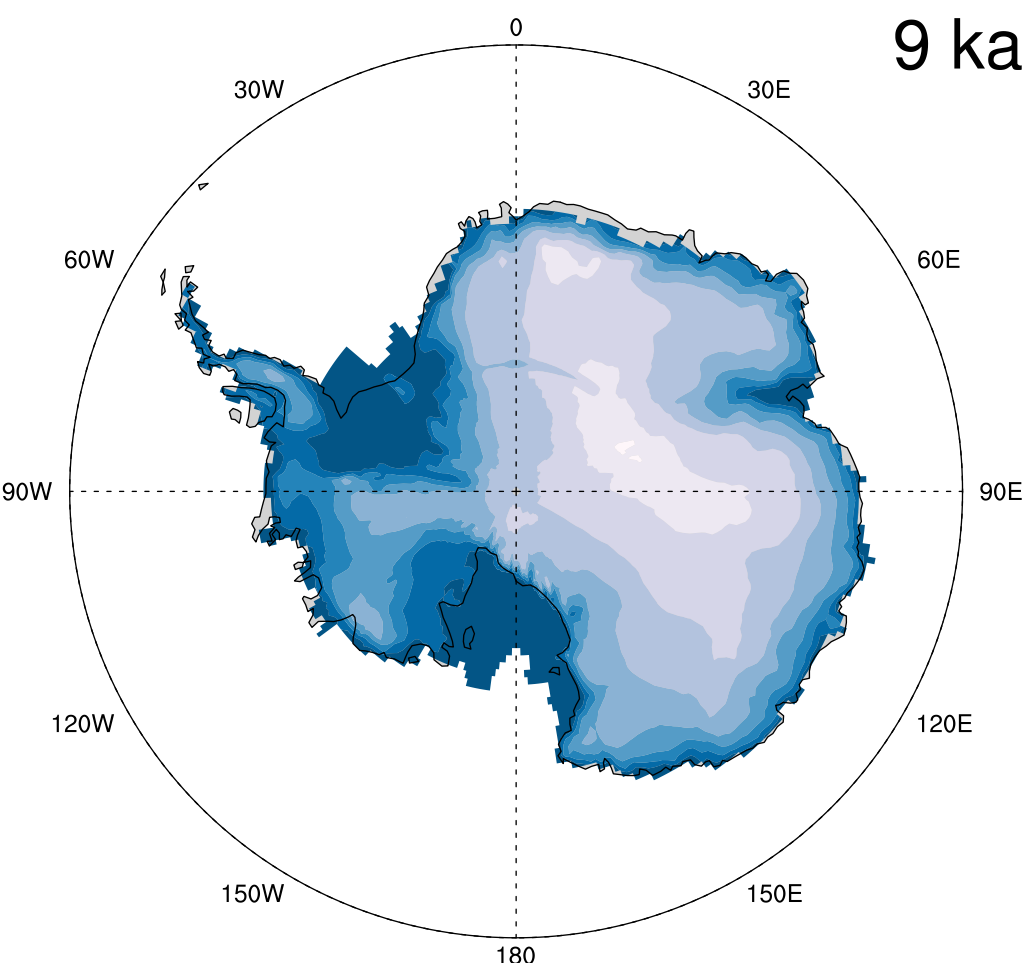
21 ka



12 ka



9 ka



0 500 1000 1500 2000 2500 3000 3500 4000

Elevation above sea level (m)