



Validation of
3D-CMCC Forest
Ecosystem Model
(v.5.1)

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Validation of 3D-CMCC Forest Ecosystem Model (v.5.1) against eddy covariance data for ten European forest sites

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Abstract

This study evaluates the performances of the new version (v.5.1) of 3D-CMCC Forest Ecosystem Model (FEM) in simulating gross primary production (GPP), against eddy covariance GPP data for ten FLUXNET forest sites across Europe. A new carbon allocation module, coupled with new both phenological and autotrophic respiration schemes, was implemented in this new version. Model ability in reproducing timing and magnitude of daily and monthly GPP fluctuations is validated at intra-annual and inter-annual scale, including extreme anomalous seasons. With the purpose to test the 3D-CMCC FEM applicability over Europe without a site-related calibration, the model has been deliberately parameterized with a single set of species-specific parameterizations for each forest ecosystem. The model consistently reproduces both in timing and in magnitude daily and monthly GPP variability across all sites, with the exception of the two Mediterranean sites. We find that 3D-CMCC FEM tends to better simulate the timing of inter-annual anomalies than their magnitude within measurements uncertainty. In six of eight sites where data were available the model well reproduces the 2003 summer drought event. Finally, for three sites we evaluate if a more accurate representation of forest structural characteristics (i.e. cohorts, forest layers) and species composition can improve model results. In two of the three sites results reveal that model slightly increases its performances, although, statistically speaking, not in a relevant way.

1 Introduction

Terrestrial ecosystems have a relevant role in the global carbon cycle, acting also as climate regulators (Peters et al., 2007; Bonan, 2008; Huntingford et al., 2009). In fact terrestrial ecosystems store large carbon stocks and cause most of the variance of carbon exchange between the atmosphere and land surfaces (Batlle Bayer et al., 2012). Among terrestrial ecosystems, forests are an essential component in the global carbon

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cycle because of their high capacity to store carbon in the vegetation and soil pools (Kramer et al., 2002). Through Gross Primary Production (GPP) plants fix atmospheric carbon dioxide (CO_2) as organic compounds, enabling terrestrial ecosystems to offset part of the anthropogenic CO_2 emissions (Janssens et al., 2003; Cox and Jones, 2008; Battin et al., 2009). Consequently, changes in GPP could have relevant impacts on atmospheric CO_2 concentration. Thus, accurately simulating terrestrial GPP is key to quantifying the global carbon cycle and predicting the future trajectories of the atmospheric CO_2 concentration (Wu et al., 2015), and taking into account the various spatial and temporal scales of the processes is a major challenge (Yuan et al., 2007).

Terrestrial ecosystem models, used to simulate carbon, water and energy fluxes, are valuable tools for advancing the knowledge of the role of ecosystems in maintaining a multitude of their fundamental services, like the provision of products and the regulation of climate (Ibrom et al., 2006). Such numerical models are also useful to: (1) predict the impacts of climate variability on terrestrial biosphere and related carbon fluxes (Ciais et al., 2005; Brèda et al., 2006; Richardson et al., 2007), ranging from long term anomalies (Santini et al., 2014) up to extreme events (Zscheischler et al., 2014); and (2) reproduce biophysical and biogeochemical feedbacks of vegetation cover and change on climate, especially when coupled to atmosphere–ocean climate models through land surface schemes (Bonan, 2008; Arneth et al., 2012; Taylor et al., 2012).

At European level, terrestrial ecosystems have been reported to be a significant sink of CO_2 (Luyssaert et al., 2012), with forests playing a relevant role in absorbing anthropogenic emissions for about 10 % (Nabuurs et al., 2003; FOREST EUROPE et al., 2011).

In the last decades some studies have identified systematic errors when modelling terrestrial ecosystem sensitivity to climate variability at multiple time scales (Friedlingstein et al., 2006; Piao et al., 2013; Dalmonech et al., 2015) while sometimes differences in model predictions are stubbornly large (F. Wang et al., 2014).

To improve the models capacity in reproducing relevant processes related to the land carbon cycle, detailed representation of missing processes should be increasingly de-

To investigate these issues, we introduced a 3-D canopy representation into the 3D-CMCC FEM, while otherwise maintaining its flexibility and the generic features to be applied to different forest ecosystems. The new model can now run on a daily time step and includes an improved allocation-phenology scheme (with the non-structural carbon pool, NSC), and an improved computation of autotrophic respiration.

2 Materials and methods

2.1 Model description

The three-dimensional Forest Ecosystem Model, 3D-CMCC FEM (Collalti, 2011; Collalti et al., 2014) (executable is available upon request at <http://www.cmcc.it/models/3d-cmcc-fem-three-dimension-forest-ecosystem-model>) is hybrid between an empirical and a process-based model relying on the concepts of the LUE approach at canopy level for carbon fixation (see Appendix A for a detailed description of algorithms). The 3D-CMCC FEM is designed to simulate at hectare scale and on a daily time step tree growth as well as carbon and water fluxes, at species level, representing ecophysiological processes in heterogeneous forest ecosystems including complex canopy structures. The 3D-CMCC FEM uses daily meteorological data, site-specific data and ecophysiological data (e.g. maximum canopy conductance, specific leaf area, etc.; see Collalti et al., 2014) to simulate forest processes. The model code architecture allows aggregating trees into representative classes, each characterized with its variables (e.g. carbon pools, leaf area index, tree height) based on their ages, species-specific and structural traits that are identified by the model through four indexes: i.e. species (x index), diameter class (Diameter at Breast Height, DBH) (y index), height class (z index), and age cohort (k index); such indexes represent the main state variables considered by the model in distinguishing ecosystems across sites. To deal with forest heterogeneity within and across different ecosystem, 3D-CMCC FEM uses a species-specific parameterization for each species simulated. Moreover, based on the assump-

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site) using only literature data (e.g. Breuer et al., 2003; Mollicone et al., 2003; Pietsch et al., 2005; White et al., 2000) independently from site-related measurements (for a full list of model ecophysiological and structural species-related parameters see Collalti et al., 2014). As in Naudts et al. (2014), in case of multiple values for a single parameter the mean values were used. Using the mean parameter estimates avoided hidden model-tuning and largely reduces the likelihood that simulation results are biased by hidden calibration.

In addition, several studies (Bolstad et al., 1999; Griffin et al., 2001; Ibrom et al., 2006; Misson et al., 2007; Cescatti et al., 2012; Guidolotti et al., 2013; Migliavacca et al., 2015) claim that beside environmental variables, spatial heterogeneity (horizontal and vertical) of the stand structure and composition (age, species) also plays an important role at the ecosystem level.

To evaluate if a more detailed simulation of forest heterogeneity improves model performances, a number of replicated simulations were performed for three heterogeneous sites (BE-Bra, IT-Ren and DE-Tha), based on different model initializations in terms of forest layers, species composition and/or ages (Table 1). These replicates start from a forest representation very close to reality (e.g. cohorts, mixed species composition and different canopy layers) to a more generalized one. For reasons of comparability, in these test sites the model has been forced with the same meteorological input data, and eco-physiological species-related parameterizations, i.e. only model initializations data, related to stand attributes, differ. These data are based on different sources: site measurements and/or literature data and/or experimental settings.

In the case of BE-Bra we initialized the model with near all possible combinations of initialization datasets. The first simulation (BE-Bra P_Q-3L) has explicitly taken into account the site heterogeneity (vertical and horizontal) (following Gielen et al., 2013, and ancillary data sources) consisting in mixed species composition at a different canopy coverage rate of *Quercus robur* (Q) and *Pinus sylvestris* (P) (20 and 80 %, respectively), with two cohorts (oaks and pines, 65 and 72 years old, respectively) and three forest layers. In the second simulation (BE-Bra P) only single-layer of Scots pines was

sites. NRMSE was quantified as:

$$\text{NRMSE}_{\text{GPP}} = \frac{\sqrt{\frac{1}{N} \sum_{i=1}^N (\text{GPP}_{\text{EC}_i} - \text{GPP}_{\text{MD}_i})^2}}{\sigma(\text{GPP}_{\text{EC}_i})} \quad (1)$$

where i represents the day (or month), and $\sigma(\text{GPP}_{\text{EC}})$ is the standard deviation of the full daily (or monthly) series of observed GPP consisting of N records.

In addition, model performances were measured for the same series through the “Model Efficiency” index (MEF) following Reichstein et al. (2002) and Migliavacca et al. (2015):

$$\text{MEF} = 1 - \frac{\sum_{i=1}^N (\text{GPP}_{\text{EC}_i} - \text{GPP}_{\text{MD}_i})^2}{\sum_{i=1}^N (\text{GPP}_{\text{EC}_i} - \text{avg}(\text{GPP})_{\text{EC}})^2} \quad (2)$$

In contrast to correlation coefficient r , the MEF index (Bowman and Azzalini, 1997) measures not only the correlation between modelled and observed data (in other words, how well they reproduce the phase of observations), but also their “coincidence”, i.e. the deviation from the 1 : 1 line, and it is sensitive to systematic deviations between model and observations (Reichstein et al., 2002).

An additional index was the Bias (Bi):

$$\text{Bi} = \frac{1}{N} \sum_{i=1}^N (\text{GPP}_{\text{MD}_i} - \text{GPP}_{\text{EC}_i}) \quad (3)$$

calculated at both annual and seasonal level, positive biases indicate an overestimation and negative values indicate an underestimation, respectively, by the simulation (see Balzarolo et al., 2014).

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To evaluate the model performances in terms of variability patterns, we adopted a procedure to compare each GPP_{EC} value to both its correspondent GPP_{MD} value and the $GPP_{EC}-GPP_{MD}$ difference, at daily and monthly level. Since the different sites have different ranges of GPP, we grouped time series values into 18 clusters, with a 5 percentile criteria, from the 5th to the 95th (Vetter et al., 2008), and we calculated the median for each group.

In order to assess the Inter-Monthly and Inter-Annual Variability (IMV and IAV respectively), individual GPP values for each month and year considered were normalized following Vetter et al. (2008) and Keenan et al. (2012). Shortly, we subtracted the respective observed or modelled average from individual (monthly and yearly) observed and modelled value as follows:

$$IMV_{(EC \text{ or } MD)_i} \text{ or } IAV_{(EC \text{ or } MD)_i} = GPP_{(EC \text{ or } MD)_i} - avg(GPP)_{(EC \text{ or } MD)} \quad (4)$$

where $avg(GPP)$ is the long-term (full series of all the available years) average of monthly (for IMV) or yearly (for IAV) GPP from observations (EC) and modeling (MD), respectively. A kernel density estimation (kde) was performed to qualitatively observe probability distribution functions (PDFs) respectively of the IMV and IAV values (Bowman and Azzalini, 1997).

To evaluate 3D-CMCC FEM ability in reproducing the observed IMV and IAV, we calculated the NRMSE based on monthly and annual time series of IMV and IAV values, respectively. The NRMSE, adopted as a normalized index of error allowing comparability among different sites, was thus calculated as in Eq. (1) but using IMV and IAV instead of GPP individual values, following the approach of Keenan et al. (2012).

on a daily scale an average correlation of 0.72 and 0.77 respectively, while on monthly scale a correlation of 0.82 and 0.86 with two low values of 0.05 and 0.06 for monthly DJF and MAM for IT-Cpz.

Winter and summer monthly average NRMSE of 1.19 and 0.97, respectively, were not significantly different to the 0.67 and 0.58 of spring and fall. MEF and Bi indexes values suggest similar findings than NRMSE.

Figure 6 shows overall modelled vs. observed fluxes over daily and monthly scales, and the absolute difference (GPP_{MD} minus GPP_{EC}) vs. observed fluxes (GPP_{EC}) as calculated by the difference matrix described in Sect. 2.4. Overall, the aggregated data reveal high correlation also due to a progressively reduced range of data, and then variability, at higher GPP values (Fig. 6a and b). Figures 6c and d show patterns of absolute difference between GPP_{MD} and GPP_{EC} with increasing GPP_{EC} . These differences result in strong reduction of discrepancies for GPP_{EC} greater than $8.5 \text{ g C m}^{-2} \text{ d}^{-1}$ for daily, or $7.3 \text{ g C m}^{-2} \text{ d}^{-1}$ for monthly temporal series (data extracted from Fig. 6c and d).

The average intra-annual GPP variations are analysed by calculating the long-term average and standard deviation values for each month of the year (Fig. 7). In spring the model results from deciduous forests present a larger variability than the observed data, especially during budburst and in late spring. The model generally matches the observed phenology timing (budburst, peak LAI, leaf senescence and their fall, i.e. length of growing season, data not shown). Consistent biases were observed in late summer.

3.2 Inter-monthly and inter-annual variability

The distribution of the IMV for the analysed sites reveals in general lower variance for modelled than observed data (Fig. 8, Table 2). Regarding deciduous forests, both DK-Sor and FR-Hes show IMV_{MD} distributions with larger interquartile range in comparison with IMV_{EC} (p value < 0.05). Conversely, for DE-Hai and IT-Col the IMV_{MD} variance is statistically representative for the IMV_{EC} ; however IT-Col shows a significantly biased

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median (p value < 0.05). Less variability than IMV_{EC} is generally observed for IMV_{MD} of conifers. While DE-Tha shows significant agreement for both variance and central tendency (p value < 0.05), at FI-Hyy the IMV_{MD} appears statistically in disagreement with IMV_{EC} for both variance and central mean tendency (Table 2). We find a small difference between IMV_{MD} and IMV_{EC} probability density modal values in IT-Ren (Table 2). Concerning broadleaved evergreen vegetation, we observe very good agreement between observed and modelled IMV central tendency measures in FR-Pue with most of the frequencies between $\pm 2gCm^{-2}d^{-1}$. In FR-Pue, however, we notice that the distributions are slightly shifted, especially around the median, with resulted variance from modelled data in disagreement with that from observed data. We detect high IMV distributions disagreement in IT-Cpz, where the PDF from observed IMV is normally distributed, while the one from modelled IMV is not (as resulted by a χ^2 goodness of fit test). IMV_{MD} series in BE-Bra (mixed forest) are in low agreement with those from EC. Modelled variance is low, and especially positive IMV values are scarcely represented. Table 2 also shows the NRMSE for IAV and IMV series. There is apparent correlation neither between sites species and average error, nor between distributions uniformity and NRMSE. In fact the lowest NRMSE for IMV was found in BE-Bra and IT-Col, the highest in DE-Hai and DK-Sor. On average the model has a NRMSE for IMVs of about 1.2.

Figure 9 shows the modelled and measured individual IAV values for each studied site. The magnitude of IAV_{MD} was on average of the same order than IAV_{EC} , showing the model ability to reproduce the inter-annual variability range, and capturing about 62 % of the anomalies signs (i.e. timing) for the total set of years. The model generally better captured conifers' IAV sign (i.e. DE-Tha, FI-Hyy, and IT-Ren), 66 % of the times against about 59 % for the deciduous forests (i.e. DE-Hai, DK-Sor, FR-Hes, IT-Col) and 55 % for the Mediterranean ones (i.e. FR-Pue and IT-Cpz). However the IAV difference in magnitude was better represented for deciduous forests rather than conifers, as inferred by the average NRMSE of respectively 1.45 and 1.66 (calculated by averaging values reported in Table 2). Although model reproduced well the timing of anomalies in

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phenology behaviour of modelled species does not cause a marked difference in the seasonal GPP cycle. Differently, IT-Ren initialized as a single layer and with one single cohort (IT-Ren 1L_1C) instead of two layers and two cohorts (IT-Ren 2L_2C) differs strongly from observed GPP values overestimating for 43.2% the annual cumulated GPP. However, for this site, the analysis of performance indices based on daily and monthly series shows no evidence of improved model results.

4 Discussions

In this paper we have analyzed the capability of the latest version of the 3D-CMCC FEM to simulate intra-annual to inter-annual GPP variability over ten heterogeneous European forest sites representative of different ecosystems and bioclimatic regions by comparing model results with observations based on EC technique. Although the model provides a reasonable reproduction of the observed values, we may evince some critical issues. First, the observed GPP data are affected by high uncertainties (Kenan et al., 2002; Papale et al., 2006; Richardson et al., 2012a, b). According to Luyssaert et al. (2007) these uncertainties in the ten case studies here considered, although at the biome level, have a very high spread, varying from ± 557.9 (for FI-Hyy) to $\pm 700 \text{ g C m}^{-2} \text{ yr}^{-1}$ (for IT-Cpz). Beside uncertainty in EC technique, model assumptions and parameterizations can increase discrepancies compared to observed GPP data.

A potential further source of error in the model runs that may need to be considered or accounted for is related to our choice of not making a site-specific parameterization. Since we used general parameterizations, large uncertainties could be detected especially in the variables that determine for example the length of the growing season (Richardson et al., 2010), and the latitudinal differences (acclimation) of the maximum, minimum and optimum temperatures for photosynthesis. Improvement could be achieved with a site-specific parameterization, but this falls beyond our goal to make the model generally applicable.

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On average, 10 years of simulations for each site have been conducted. In addition, in three sites different model initializations (i.e. considering different forest structure, composition and cohorts) were used to quantify improvements in model results when a more detailed heterogeneity forest structure representation and processes are simulated. Modelled GPP results were compared against those from EC observations collected for these sites encompassing three mono-specific (pure) stands of Beech, Holm oak and Scots pine, and three uneven-aged, multi-layered and mixed stands.

Based on results, we can now provide answers to the four initial questions.

4.1 Does the model reproduce the magnitude and timing of seasonal fluctuations in GPP across different forest types, structures and compositions?

Overall, as desirable, the model is skilful in reproducing the annual cumulated and intra-annual (seasonal) cycle of GPP, calculated as both daily and monthly value average, with the monthly scale performing better across all statistical indices considered. These results can be anyway considered as a “false positive” due to the strong seasonality of GPP patterns that influences and causes high values of correlation more than model capabilities to reproduce GPP fluxes. This is clearly related to the tendency to linearize the relationship between CO_2 flux and PAR, as also reported by Ruimy et al. (1995) and Wu et al. (2015), respectively. Overall, average annual modelled values were highly consistent with EC data, except for the Mediterranean sites (where seasonality is less pronounced). Here summer drought stress showed to be the most limiting factor on photosynthesis at FR-Pue (Falge et al., 2002; Reichstein et al., 2002; Sabatè et al., 2002) while the presence of shallow groundwater table at IT-Cpz seems reducing the severity of summer drought. However the model showed non-negligible uncertainties in representing GPP patterns, as inferred by temporal mismatches in variance. The overall agreement despite temporal mismatches suggested that errors compensated over the year, but are cumulated in specific time windows (e.g. seasons). As reported for other models (Morales et al., 2005; Naudts et al., 2014) the model’s performances

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are generally worse in winter (DJF) and summer (JJA). Biases and differences in winter GPP variance may be related to the model algorithms used to simulate LAI and to the algorithm used to calculate GPP from EC data (Reichstein et al., 2005), since GPP variability should be low during DJF, especially for deciduous forests. However, mismatches are also related to the way in which 3D-CMCC FEM represents winter and early spring ecosystem processes. The model in fact does not consider the influence of ground vegetation that appears to be not negligible in some cases (Kolari et al., 2006).

High GPP variance for evergreen species could be strongly related to low temperatures during winter (Delpierre et al., 2009). Systematic overestimation in winter and spring GPP could then be associated with lacks in representing conifers acclimation or to soil and atmosphere thermal constraints. At high latitudes and altitudes another source of uncertainty may be related to freezing and thawing dynamics in soil water (Beer et al., 2007) which are not considered by the model, as like as snow sublimation and melting are still simplistically represented.

GPP of deciduous forests in summer and autumn are also affected by uncertainties for surface, which is represented by LAI in the model. In addition GPP is linear with respect to PAR (Monteith and Moss, 1977) over monthly or annual time scale, while the relation is strongly nonlinear at daily scale (Leuning et al., 1995; Gu et al., 2002; Turner et al., 2003; Wu et al., 2015). The linear response of GPP to PAR led to the underestimation/overestimation of GPP under conditions of low/high incident PAR (Propastin et al., 2012; He et al., 2013). In case of stress or photoinhibition leaves reduce or stop the photosynthesis at too high levels of radiation, while in normal condition, photosynthesis is light-saturated at high PAR (Mäkelä et al., 2008) which lets canopy photosynthesis saturated at relatively low PAR even in dense tropical forests with high LAI (Ibrom et al., 2008). The model overestimation of summer GPP may thus be partially related to the lack of representation of canopy photosynthesis saturation processes.

Although adopting a more complex phenology scheme, in the comparison between deciduous and evergreen forests our model showed better performances for deciduous

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is sparse, the phenology of understory vegetation (Richardson et al., 2010). Leaves of different ages have different efficiency, sensitivity to solar radiation, temperature and water related stresses (Chabot and Hicks, 1982). All these elements may have an important role in affecting GPP dynamics, but are still scarcely or not represented by mechanistic ecosystem or forest models. As a confirmation of these suspects, slight modifications in representing phenology and leaf turnover resulted in general improvement of model consistency with EC data (Marconi, 2014).

Distribution of IMV values showed specific patterns attributable to the dominant species. Beech forests IMV PDFs were concentrated around the average value and strongly influenced by high biases. This pattern was probably due to the fact that half of the months in one year have no or little photosynthesis (i.e. early spring, fall and winter) and most of the photosynthetic activity occurs in late spring and summer, when carbon assimilation is influenced by temperatures and solar radiation (Mercado et al., 2009). Conifers PDFs were usually smoother, non-skewed, with reduced variability and fitted by a statistical normal curve.

The model showed an average NRMSE for IMV of 1.22 but still captured about two third of the annual anomalies sign.

The results for IAV (see Fig. 9) are quite contrasting, and largely depend on site and the number of annual-by-annual comparison. Better results have been obtained for FI-Hyy and FR-Pue, so there is not apparent correlation with latitudes and forest species. Similarly, lower results are reported for IT-Ren, IT-Cpz and BE-Bra where the number of annual correlations are lower than the other sites. The magnitude of differences in the standard deviation follows generally the same tendency, particularly for BE-Bra, IT-Ren and IT-Cpz. These results confirm the model limited ability to represents the inter-annual variability in these specific sites rather than in these ecosystems. The comparison between modelled and observed data at the inter-annual time scale shows the model to be sufficiently able to reproduce the sign of variability through the years including the extreme events (heat wave combined to drought) during the 2003 summer (Ciais et al., 2005; Vetter et al., 2008) and, for some sites, the anomalous

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most of the times their magnitude was not. This behaviour seems to be related to several aspects, mainly to an over/under estimation of the causes that reproduce anomalies, e.g. processes simulated linked to the type of climate anomaly, mismatches in phenology or to a missed representation of others processes (e.g. mast years, disturbances, shallow water). Keenan et al. (2012) asserts that lacks in phenological variability and in canopy and soil dynamics are the main culprits of these mismatches but also that flux measurements are affected by random errors especially when fluxes are higher. Poulter et al. (2009) founded similar magnitude of errors also with models that were driven by remote-sensing data. Open questions remain as to the proportion of interannual variability in land–atmosphere carbon exchange that is directly explainable by variability in climate (Hui et al., 2003; Richardson et al., 2007).

4.3 Is the model generic enough that a single set of species-specific parameterization allows reproducing GPP behaviour across different ecosystems without further need of a site-related calibration?

Overall the model showed good flexibility although the sites showed a pronounced spatial and temporal heterogeneity (i.e. a variable number of forest layers, different cohorts and species). The model was able to reliably represent the ecophysiology of beech and spruce species at different latitudes, without modifying or tuning the parameterization sets. However, annual and seasonal performance indices, calculated exploiting daily and monthly series, evidenced different performances between the two northern beech sites and the two southern ones. Tables S1 and S2 show a systematic difference in all the statistics used, suggesting the presence of a latitudinal gradient in 3D-CMCC FEM ability to represent beech forest processes. This gradient could be explained by how the model represents the different limiting factors and their impacts on GPP. For example we expect low temperatures to be the most important limiting factor at higher latitudes, whereas soil water availability at lower latitudes (Chapin et al., 2002).

We had similar results for the two spruce sites. The model showed better performance at higher latitudes. While phenotypic plasticity, and thus the parameter set, may

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still remain especially during winter and summer. The first reason for these low agree-
ments in winter can be attributable to errors during the estimate of GPP from NEE and
Ecosystem Respiration values from measurements data. The second can be related to
the model's lack or simplicity in representation of snow pack dynamics as reported by
Krishnan et al. (2008, 2009) especially for evergreen sites (Keenan et al., 2012). Dis-
agreements in summer could be related to model simplicity in simulating soil drought
and, using the Monteith approach (Monteith and Moss, 1977), to the strong nonlinearity
at daily scale of GPP and PAR, and to the lack of representation of the light saturation
processes. In addition, as reported by Keenan et al. (2012), the apparent high variabil-
ity in the data during the summer season could therefore be due to random errors in
the flux measurements generating larger variability and then lower correlations against
modelled data.

No marked differences were found in simulations across different latitudes, so model
parameterizations for the different tree species could be useful over Europe with a quite
high rate of confidence, with the exception of specific cases in Mediterranean forests.

As for other models, 3D-CMCC FEM showed to have the potential to correctly repro-
duce the signs of interannual variability, like the 2003 heat wave and drought extreme
and the anomalous carbon uptake during the warm spring of 2007 and their instan-
taneous biological response to these events. Significant disagreements were however
found in reproducing magnitude of these anomalies.

The consideration of stand heterogeneity, when possible or existing (i.e. layers,
cohorts and mixed composition), led the model to improve its results in two of the three
sites compared to generalized simulations of forest attributes. This plasticity makes the
model able to be used in a wider range of forest ecosystems.

Appendix A: Model description

A1 Photosynthesis

As in the Collalti et al. (2014) 3D-CMCC FEM version, the carbon flux is still estimated by multiplying, for a particular species x , the absorbed photosynthetic active radiation (APAR, i.e. the radiation intercepted by the canopy) with the leaf area index (LAI, $\text{m}^2 \text{m}^{-2}$) with either the prognostic potential radiation use efficiency (ε_x , grams of dry matter MJ^{-1}) or the maximum canopy quantum use efficiency (α_x , $\mu\text{molCO}_2 \mu\text{mol}^{-1}$ PAR) (for a full list of model parameters, algorithms, and indexes see Collalti et al., 2014). Parameters ε_x or α_x are controlled by the product of several environmental factors (modifiers) indicated as $\text{mod}_{x,k}$ (dimensionless values varying between 0 and 1 and differing for each species x and age class k) depending on: vapour pressure deficit, daily maximum and minimum air temperatures, soil water content and site nutrient status (for a full modifiers description see Landsberg and Waring, 1997). Gross primary production (GPP; $\text{gC m}^{-2} \text{day}^{-1}$) is thus calculated using the following equation:

$$\text{GPP}_{x,y,z,k} = \varepsilon_x \cdot \text{APAR}_z \cdot \text{mod}_{x,k} \quad (\text{A1})$$

where APAR is the absorbed radiation by the trees at the z th layer (where z represents the layer of representative height for each height class), while y represents the tree diameter class.

Autotrophic Respiration (AR) is treated distinguishing into Maintenance Respiration (MR), governed by a Q_{10} type response function (see Sect. A4) (Ryan, 1991; Bond-Lamberty et al., 2005) and Growth Respiration (GR) assumed to be a constant proportion (30%) of all new tissues produced (Larcher, 2003). Net Primary Production (NPP), is calculated as follows:

$$\text{NPP}_{x,y,z,k} = \text{GPP}_{x,y,z,k} - \text{AR}_{x,y,z,k} \quad (\text{A2})$$

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NPP is then partitioned into biomass compartments and litter production following dynamic allocation patterns that reflect environmental constraints (i.e. light and water competition) and age.

A2 Daily meteorological forcing and snow dynamics

5 The model implements a daily time step due to the temporal frequency of meteorological forcing input data; average maximum (T_{\max}) and minimum air temperature (T_{\min}), soil temperature (T_{soil}), vapour pressure deficit, global solar radiation and precipitation. In addition, the model uses the day-time (T_{day}) and night-time (T_{night}) average temperature computed as follows (Running and Coughlan, 1988):

$$10 \quad T_{\text{day}} = 0.45 \cdot (T_{\max} - T_{\text{avg}}) + T_{\text{avg}} \quad (\text{A3})$$

$$T_{\text{night}} = (T_{\text{day}} + T_{\min})/2 \quad (\text{A4})$$

When the soil temperature is missing among in situ observed data, the model estimates it for the upper 10 cm of the soil layer through an 11 day running weighted average of daily average air temperature and further corrected by the presence of a snowpack as in Thornton (2010), Kimball et al. (1997) and Zeng et al. (1993). The variable related to the snowpack thickness was included as a water cycle component by reproducing the daily amount (mm day^{-1}) of snow melt driven by average air temperature (T_{avg}) and incident net global radiation (Rad_{soil}), while snow sublimation is only driven by average air temperature.

20 In case of snow presence, if the average air temperature is higher than 0°C , considered the melting point as in Running and Coughlan (1988) and Marks et al. (1992), the rate of daily snowmelt is estimated by:

$$\text{Snow}_{\text{melt}} = (t_{\text{coeff}} \cdot T_{\text{avg}}) + \left(\frac{\text{Rad}_{\text{soil}} \cdot \varepsilon_{\text{snow}}}{H_{\text{fus}}} \right) \quad (\text{A5})$$

where t_{coeff} is the snowmelt coefficient ($0.65 \text{ kg m}^{-2} \text{ }^{\circ}\text{C}^{-1} \text{ day}^{-1}$), ϵ_{snow} is the absorptivity of snow (0.6), H_{fus} is the latent heat of fusion (335 kJ kg^{-1}), Rad_{soil} is the incident net global radiation at the soil surface ($\text{kJ m}^{-2} \text{ day}^{-1}$).

Otherwise, if the average air temperature is lower than 0°C snow sublimation is computed by:

$$\text{Snow}_{\text{subl}} = \left(\frac{\text{Rad}_{\text{soil}} \cdot \epsilon_{\text{snow}}}{H_{\text{sub}}} \right) \quad (\text{A6})$$

where H_{sub} is the latent heat of sublimation (2845 kJ kg^{-1}).

A3 Phenology and carbon/nitrogen allocation

Phenology plays a fundamental role in regulating photosynthesis and other ecosystem processes (e.g. carbon and nitrogen dynamics), as well as inter-individual and inter-species competitive relations and feedbacks to the climate system (Richardson et al., 2012a). In the updated model version phenology and carbon allocation depend on six different carbon and nitrogen pools. Five pools represent the main tree organs: foliage, (fine and coarse) roots, stem, branch and bark fraction. One pool corresponds to non-structural carbon (starch and sugar) stored in the whole tree. Woody pools are further distinguished between live and dead wood. This is necessary to represent NSC mobilization and consequently leaf phenology (e.g. leaf production during spring for deciduous trees) and carbon allocation. In the new version of 3D-CMCC FEM LAI values are predicted for sun and shaded leaves (De Pury and Farquhar, 1997; Thornton and Zimmermann, 2007; Wu et al., 2015), minimizing the effects of the “Big-leaf” approach (Monteith, 1965; Sellers et al., 1997), as a function of the amount of carbon allocated to the leaf pool. It is noteworthy that each pool and each structural state variables is daily updated according to the meteorological data, forest structure and simulated fluxes. Following Arora and Boer (2005), for deciduous species the model considers five phenological transitions that drive the seasonal progression of vegetation through phases

of dormancy/quiescence, budburst, maximum growth, active growth, and senescence as in the following:

1. Leaf onset starts from quiescence when thermic sum (the sum of the T_{day} air temperatures exceeding the threshold T_{base} value of 5°C) exceeds a species- and site-specific temperature threshold value (Rötzer et al., 2004; Dufrene et al., 2005) and up to $\text{LAI} = \max(\text{LAI}) \cdot 0.5$. The costs of expanding buds during this period of high carbon demand are supported by NSC (Landhausser, 2010; Dickmann and Kozlowski, 1970).
2. During the budburst phase, carbon and NSC are allocated to the foliage pool, as long as the balance between GPP and AR is positive (Barbaroux and Bréda, 2002; Campioli et al., 2013; Scartazza et al., 2013).
3. During the succeeding maximum growth phase and lasting up to peak LAI, carbon is allocated into foliage and fine root pools (Sabatè et al., 2002), based on the pipe model theory (Shinozaki et al., 1964a, b), to optimize photosynthesis; otherwise, no growth occurs and NSC is used.
4. Successively, the full growing phase lasts up to the day when day length (in hours) is shorter than a species-specific threshold value. In this phase carbon is allocated into stem, fine and coarse roots, branch and bark, and into non-structural carbon pools in order to refill the reserves for the next years.
5. Finally, during the leaf fall (i.e. yellowing or senescence) phase, lasting until the leaf fall (assumed linear) is complete, the total positive carbon balance is allocated to the NSC pool.

Outside the growing season (dormancy) trees consume NSC for fuelling maintenance respiration (Ogren, 2000).

For evergreen species the model follows a similar but simplified approach simulating a first maximum growth phase, when the model allocates NSC to foliage and fine roots

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up to reach peak LAI, and a second full growing phase, when the model allocates to the other pools. As in Lawrence et al. (2011) for litterfall we assume and simplify that there are no distinct periods, but rather a continuous shedding of foliage and fine roots of the previous years.

All tree pools are updated at a daily time step depending on NPP. Nitrogen concentration for each pool is considered as a C/N ratio following Thornton (2010) and Dufrene et al. (2005). The C/N stoichiometry is constant and depends on species, unfortunately, the model still lacks of an interactive C-N cycle. Forest stand structural attributes, e.g. diameter at breast height (DBH), tree height, and crown competition are also updated at a daily timestep based on species-specific biometric relationships.

A4 Autotrophic respiration

Based on the approach of BIOME-BGC model (Thornton, 2010) 3D-CMCC FEM computes the daily AR of all living tissues. MR is a modified Van't Hoff function (Davidson et al., 2006; Mahecha et al., 2010) of temperature with the temperature sensitivity parameter Q_{10} (see below) and a linear function of the nitrogen content ($N_{\text{content}} = 0.218 \text{ kgC kgN}^{-1} \text{ day}^{-1}$; Ryan, 1991) in the living compartments. The Q_{10} function is an exponential function for which a 10°C increase in temperature relates to a Q_{10} factor change in the rate of respiration. MR is partitioned into day time and night time respiration using, in place of temp in Eq. (A7): t_{day} and t_{night} for foliage, t_{soil} for fine and live coarse roots, and t_{avg} for live stem and branch.

$$MR_{x,y,z,k} = 0.218 \cdot N_{\text{content},x,y,z,k} \cdot Q_{10}^{(\text{temp}-20)/10} \quad (\text{A7})$$

$GR_{x,y,z,k}$ is considered as a fixed ratio (30%) of all newly grown (i.e. living) tissues as proposed by Larcher (2003).

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Author contributions. A. Collalti conceived the paper, designed the experiments, co-developed the model code, performed the simulations, and wrote the manuscript with contribution from all co-authors, S. Marconi co-developed the model code, performed the simulations, and contributed to data analysis, A. Ibrom contributed to manuscript improvement, C. Trotta contributed to data analysis, A. Anav contributed to data analysis and to the manuscript improvement, E. D'Andrea contributed to data analysis, G. Matteucci contributed to manuscript improvement and data analysis, L. Montagnani contributed to manuscript improvement and data analysis, B. Gielen contributed to manuscript improvement and data analysis, I. Mammarella contributed to manuscript improvement and data analysis, T. Grünwald contributed to manuscript improvement and data analysis, A. Knohl contributed to manuscript improvement and data analysis, R. Valentini contributed to manuscript, M. Santini contributed to manuscript conceiving and data analysis improvements.

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Table 2. IMV and IAV NRMSE for the analyzed sites. Each specific IMV distribution was tested for normality goodness of fit (N = normal distribution, P = non normal distribution). A test for equivalence of central tendency and variance was performed between IMV_{MD} and IMV_{EC} values. (na) refers to the case of sites with inconsistent distributions (one normal, one not normal distributed). (*) marks refer to the acceptance of the null hypothesis that the two distributions are equivalent for the specific statistic ($\alpha = 0.05$). ECT stands for “Equivalence for Central Tendency”; EV for “Equivalence for Variance”.

		DE-Hai	DK-Sor	FR-Hes	IT-Col	FR-Pue	IT-Cpz	DE-Tha (1S)	FI-Hyy	IT-Ren (2L-2C)	BE-Bra (P_Q-3L)
NRMSE	IAVs	2.4	1.8	1.3	0.3	0.6	1.1	1.0	2.7	1.3	0.9
NRMSE	IMVs	1.7	2.7	1.1	0.6	1.1	1.2	1.1	1.2	1.0	0.5
ECT	p value	1.00* ^N	0.12* ^N	0.54* ^N	0.00 ^N	0.15* ^N	1.00* ^{na}	1.00* ^P	0.04 ^N	0.88* ^P	0.85* ^N
EV	p value	0.53* ^N	0.00 ^N	0.00 ^N	0.46* ^N	0.00 ^N	0.02 ^{na}	0.78* ^P	0.00 ^N	0.27* ^P	0.01 ^N

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Table 3. Performance statistics (r , NRMSE, MEF, Bi) are reported as derived from daily and monthly series of GPP_{EC} and GPP_{MD} values over long-term annual scale, for the different forest structure simulations. The (*) refers to p value < 0.0001 in correlation between GPP_{EC} and GPP_{MD} data. In addition, long term average of annual GPP_{MD} and GPP_{EC} values ($\text{gCm}^{-2}\text{yr}^{-1}$) for the different forest structures are shown.

Site	Model Set-up code	Daily				Monthly				Yearly	
		r	NRMSE	MEF	Bi $\text{gCm}^{-2}\text{d}^{-1}$	r	NRMSE	MEF	Bi $\text{gCm}^{-2}\text{month}^{-1}$	GPP_{MD} $\text{gCm}^{-2}\text{yr}^{-1}$	GPP_{EC} $\text{gCm}^{-2}\text{yr}^{-1}$
BE-Bra	P	0.72*	0.73	0.47	-0.7	0.86*	0.55	0.7	-9.81	1003	
	Q_3L	0.76*	0.91	0.18	0.14	0.84*	0.71	0.49	2.67	1105	
	Q_2L	0.74*	0.89	0.21	0.0	0.86*	0.74	0.45	9.38	1179	
	Q_1L	0.75*	0.95	0.01	0.25	0.86*	0.68	0.53	6.86	1147	1112
	P_Q-3L	0.77*	0.64	0.58	0.2	0.91*	0.42	0.82	-3.9	1169	
	P_Q-2L	0.75*	0.67	0.55	-0.41	0.91*	0.44	0.81	5.97	1037	
P_Q-1L	0.75*	0.66	0.56	-0.37	0.91*	0.68	0.53	6.86	1056		
IT-Ren	2L_2C	0.81*	0.62	0.61	-1.27	0.95*	0.3	0.91	-39.3	1348	
	1L_1C	0.83*	0.85	0.27	1.27	0.96*	0.61	0.62	38.9	1950	1362
DE-Tha	1S	0.89*	0.48	0.8	-0.1	0.96*	0.29	0.91	-2.5	1898	
	2S	0.89*	0.46	0.79	-0.08	0.95*	0.27	0.93	-2.59	1837	1869

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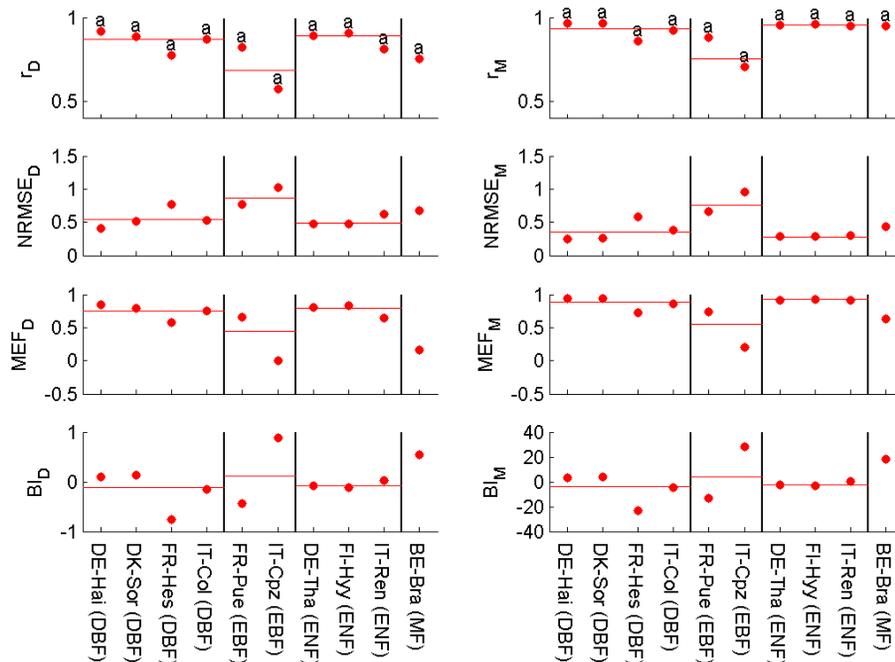


Figure 1. 3D-CMCC FEM performance indices at different time scales; daily (on the left) and daily aggregated to month (on the right). DE-Tha refers to the 1S simulation, IT-Ren to the 2L_2C simulation, BE-Bra to the P_Q-3L simulation (see text). The red horizontal line refers to the value calculated for the whole data aggregated per IGBP vegetation class. In case of r , “a” stands for p value < 0.001 .

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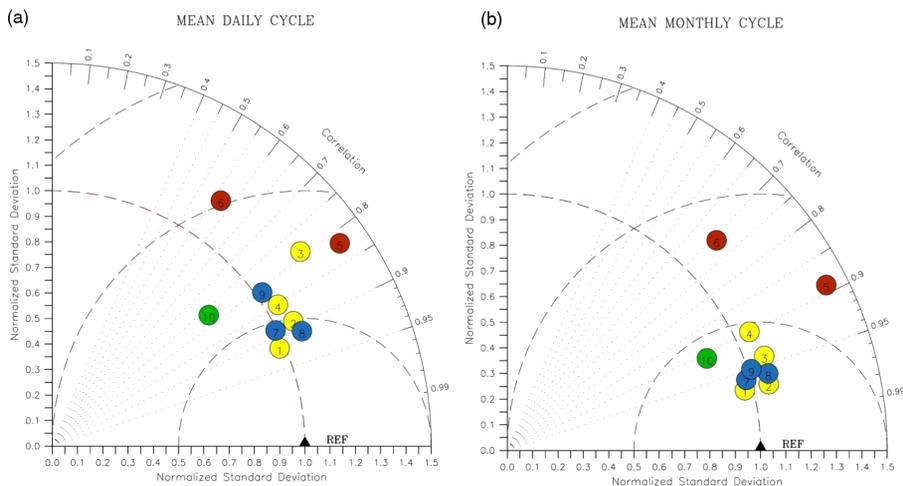


Figure 2. Taylor diagrams for daily (a), daily aggregated to month (b) GPP evaluated by: the deviation of model results from observations (REF) in terms of normalized standard deviation of observations, represented by the distance from the site point to the point on the x axis identified as reference (REF); the difference of model normalized standard deviation from that of observations, represented by the distance of the site point with respect to the quarter arc crossing REF; and the correlation, given by the azimuthal position of the site point to the x axis. The sites are numbered in ascending order as follows: Eq. (1) DE-Hai, Eq. (2) DK-Sor, Eq. (3) FR-Hes, Eq. (4) IT-Col, (5) FR-Pue, (6) IT-Cpz, (7) DE-Tha, (8) FI-Hyy, (9) IT-Ren, (10) BE-Bra. Colors refer to different IGBPs: DBF (yellow), EBF (orange), ENF (light-blue), MF (green).

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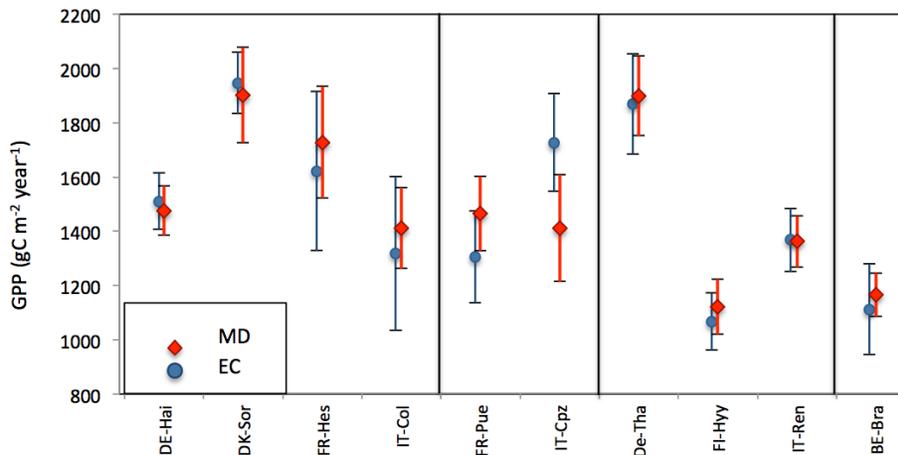


Figure 3. Distributions of annual GPP ($\text{gC m}^{-2} \text{yr}^{-1}$). MD (red) are model results, EC (blue) measured by eddy covariance. The vertical bars represent ± 1 standard deviation. DE-Tha refers to the 1S simulation, IT-Ren to the 2L_2C simulation, BE-Bra to the P_Q-3L simulation (see text).

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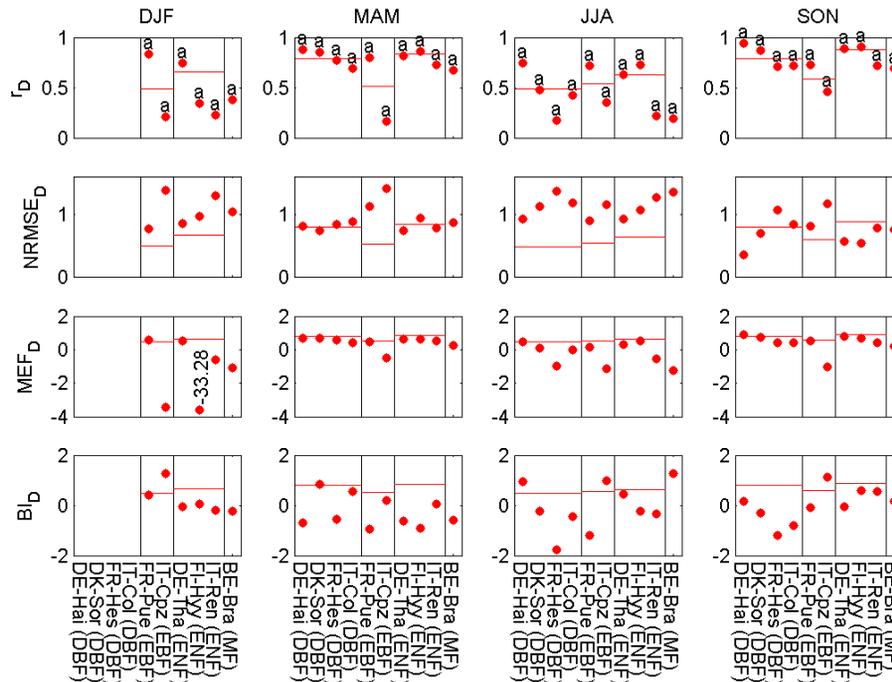


Figure 4. 3D-CMCC FEM performances indices of daily (D) GPP at different seasons. DE-Tha refers to the DE-Tha (1S) simulation, IT-Ren to the (2L-2C), BE-Bra to the (P_Q-3L). The red horizontal line refers to the value calculated for the whole data aggregated per IGBP. Strongly negative MEF are represented out of scale, but flanked with their respective numerical value. In case of r, “a” stands for $p < 0.001$.

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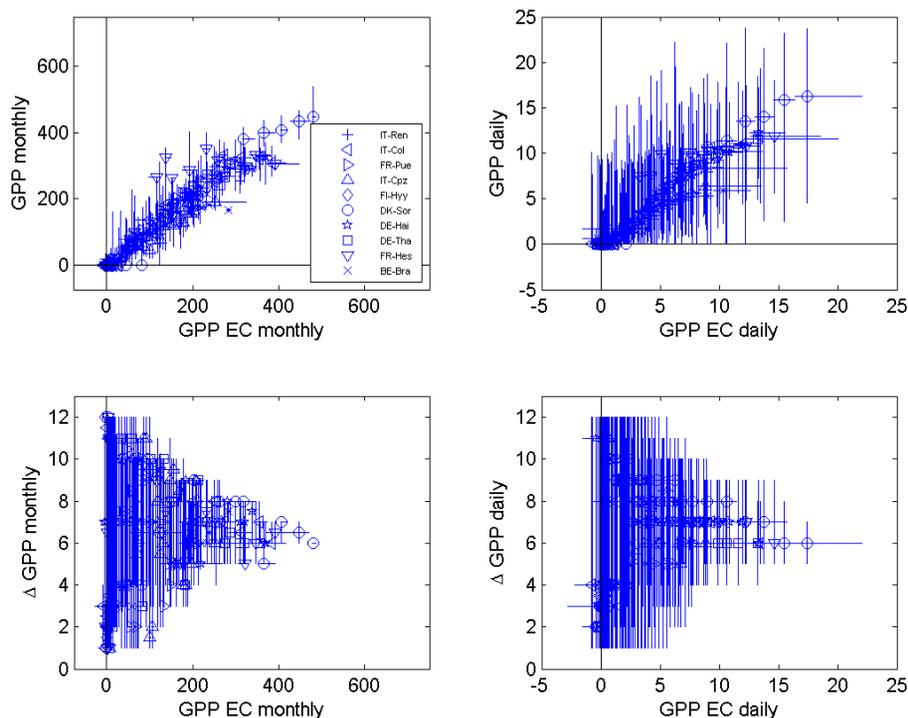


Figure 6. Comparison between GPP_{MD} and GPP_{EC} data. The top plots show the average $GPP_{EC} : GPP_{MD}$ correlation for daily ($gCm^{-2}d^{-1}$) and monthly ($gCm^{-2}month^{-1}$) data. The bottom plots show absolute difference range between GPP_{MD} and GPP_{EC} while increasing GPP_{EC} values. Negative values are excluded because of model assumptions. DE-Tha refers to the 1S simulation, IT-Ren to the 2L_2C simulation, BE-Bra to the P_Q-3L simulation.

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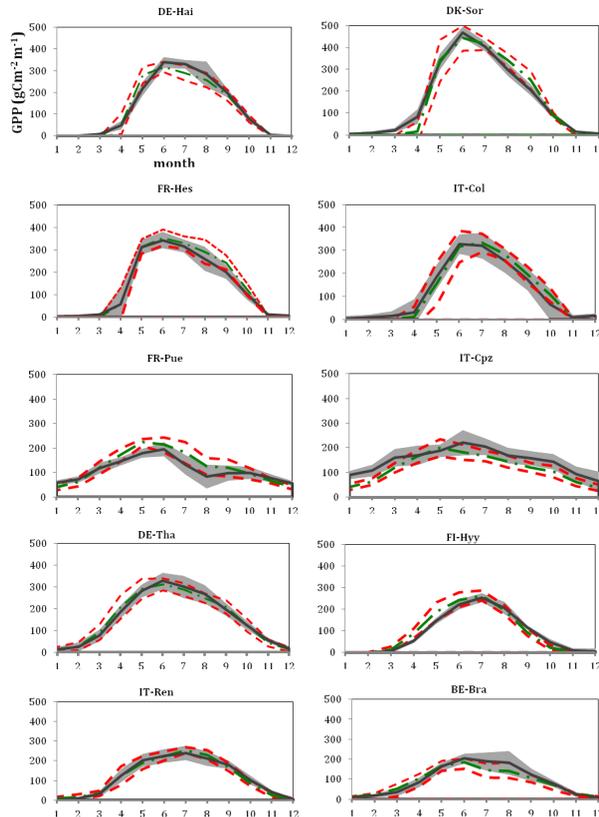


Figure 7. Seasonal (monthly) cycle of GPP across the ten sites. The grey line and margins of the grey area represent long-term average of monthly GPP_{EC} ($gC\ m^{-2}\ month^{-1}$) and its ± 1 standard deviation, respectively. The green and red dashed lines represent the long-term average of monthly GPP_{MD} ($gC\ m^{-2}\ month^{-1}$) and its ± 1 standard deviation, respectively. DE-Tha refers to the 1S simulation, IT-Ren to the 2L_2C simulation, BE-Bra to the P_Q-3L simulation (see text).

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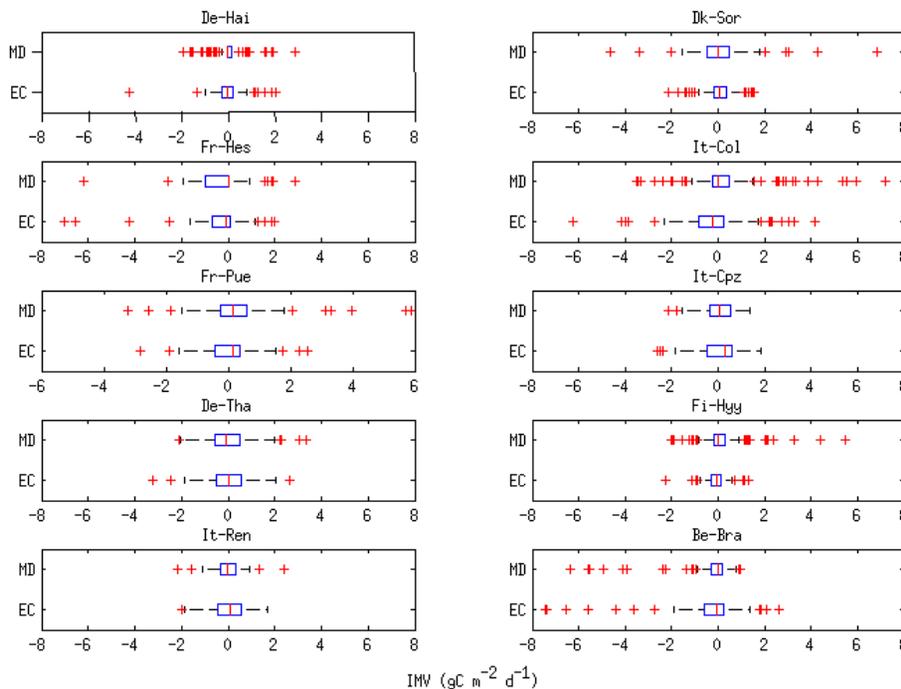
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Figure 8. Distribution of the magnitude for the inter-monthly variability values (IMVs, $\text{gC m}^{-2} \text{d}^{-1}$) for each specific site, resulted by standard kernel density estimation. DE-Tha refers to the 1S simulation, IT-Ren to the 2L_2C simulation, BE-Bra to the P_Q-3L simulation. Red vertical line represents the median, the blue box bounds the 25th and 75th percentiles, black dashed lines limit whiskers at the 0.35th and 99.65th percentiles, and red crosses are the outliers.

Validation of 3D-CMCC Forest Ecosystem Model (v.5.1)

A. Collalti et al.

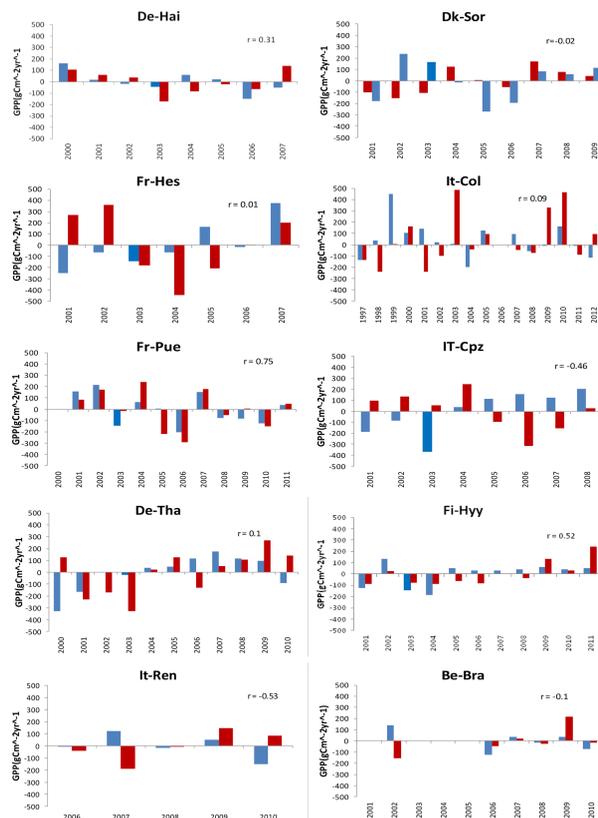
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Figure 9. Inter-Annual Variability (IAV) based on Keenan et al. (2012). Red and blue bars indicate the observed and modelled IAV values, respectively; r values refer to correlation between observed and modelled variations. DE-Tha refers to the 1S simulation, IT-Ren to the 2L_2C simulation, BE-Bra to the P_Q-3L simulation.